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Retrieval of Marine Parameters in Coastal Waters from Satellite Data (Extended Abstract)

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Abstract: A new method has been developed for simultaneous retrieval of aerosol optical properties and chlorophyll concentrations in Case 1 waters (Stamnes et al., 2003a). This method is based on an improved, complete model for the inherent optical properties (Li and Stamnes, 2002) and accurate simulations of the radiative transfer process in the coupled atmosphere-ocean system. It has been implemented in NASA's SeaWiFS Data Analysis System (SeaDAS) and tested on SeaWiFS imagery obtained off the East Coast of the US. We summarize the salient features of this approach, and describe the steps required to adapt it for application to more complex (Case 2) coastal waters.

1. Introduction

The feasibility of measuring marine chlorophyll concentrations from space was demonstrated by the CZCS (Coastal Zone Color Scanner) mission (Gordon *et al.*, 1980; Hovis *et al.*, 1980) showing that (1) there is a relationship between ocean color and chlorophyll concentration for most of the open ocean, (2) it is possible to remove the interfering effects of the atmosphere from satellite imagery. Recently sensors with more channels and higher radiometric sensitivity, such as SeaWiFS (the Sea-viewing Wide Field-view Sensor) (Hooker *et al.*, 1992), MODIS (the Moderate-Resolution Imaging Spectroradiometer) (Salomonson *et al.*, 1989), OCTS (Ocean Color and Temperature Scanner) (Saitoh *et al.*, 1995), MERIS (Medium Resolution Imaging Spectrometer) (Rast *et al.*, 1995), GLI (Global Imager onboard the Advanced Earth Observing Satellite ADEOS-II) and others, have been designed and deployed for remote sensing of ocean color.

It is customary to retrieve the chlorophyll concentration in two independent steps: (i) an atmospheric correction algorithm is applied to separate the atmospheric radiance from the radiance that comes from the ocean (the water-leaving radiance); (ii) an empirical bio-optical model is used to relate this water-leaving radiance to the chlorophyll concentration. Most bio-optical models are based on empirical or semi-empirical relations derived by statistical regression of remote sensing reflectance versus chlorophyll concentration (O'Reilly et al., 1998; Gordon et al., 1988; Reynolds et al., 2001; Carder et al., 1999; Garver and Siegel; 1997), and many of them have been validated against the SeaBAM (SeaWiFS Bio-optical Algorithm Mini-Workshop) data base

(O'Reilly et al., 1998).

Gordon and co-workers (Gordon and Wang, 1994; Wang and Gordon, 1994; Gordon, 1997) proposed an operational procedure for atmospheric correction, which has been used in the SeaWiFS algorithm. Similar procedures have been proposed for the MERIS (Antoine and Morel, 1999) and OCTS instruments (Fukushima *et al.*, 1997). These algorithms are based on the assumption that the ocean is black in the near infrared, and this black-pixel assumption is applied to the NIR channels (at 765 nm and 865 nm). Although an attempt was recently made to remedy this black-pixel assumption (Siegel *et al.*, 2000), three major sources of uncertainty still remain: (i) the decoupling of the atmospheric and oceanic radiative transfer problems makes it difficult to quantify the retrieval error, (ii) current algorithms are based on the questionable assumption that the water-leaving radiance is isotropic (Morel and Gentili, 1996; Yang and Gordon, 1997; Li and Stamnes, 2001; Yan *et al.*, 2002b), and (iii) a simple reflectance model has been used (Siegel *et al.*, 2000), in which the absorption and backscattering coefficients are empirically related to the chlorophyll concentration (Gordon *et al.*, 1975; Gordon *et al.*, 1988).

Recently we developed an atmospheric correction algorithm that does not rely on several assumptions employed in previous algorithms (Stamnes *et al.*, 2003a). This algorithm employs a radiative transfer model for the *coupled* atmosphere-ocean system based on the discrete-ordinate method (CAO-DISORT) to circumvent these unnecessary assumptions. It has been implemented in the SeaWiFS Data Analysis System (SeaDAS) and tested on ocean color data. Here we summarize the salient features of this algorithm and describe the work required to adapt it for application to coastal (Case 2) waters.

Summary of Our Case 1 Ocean Color Algorithm

To remove the uncertainties inherent in current algorithms our development is based on a rigorous discrete-ordinates solution of the radiative transfer equation pertinent to the coupled atmosphere-ocean system (CAO-DISORT (Stamnes et al., 1988; Jin and Stamnes, 1994; Thomas and Stamnes, 1999; Yan and Stamnes, 2003), which provides accurate results in a computationally efficient manner (Gjerstad et al., 2002). The CAO-DISORT code has already been used to create algorithms for retrieval of marine constituents in Case 2 waters (Frette et al., 1998; Frette et al., 2001), but these algorithms rely on a simplified description of atmospheric parameters, and have not been tested on real data. As aquatic input to the CAO-DISORT code we use an improved, complete model for the inherent optical properties (IOPs) of Case I waters that was specifically designed for this purpose, and that has been tested against remote sensing reflectances compiled in the SeaBAM data base (Li and Stamnes, 2002). We use the CAO-DISORT code to compute TOA reflectances, and store the results in lookup tables (LUTs) to enable rapid access to these reflectances $\rho^{LUT}(\lambda)$. Based on these accurate simulations of the light field throughout the coupled atmosphere-ocean system we have developed a method for simultaneous and self-consistent retrieval of atmospheric aerosol properties and chlorophyll concentrations in Case 1 waters (Stamnes et al., 2003a).

In our new atmospheric correction algorithm we employ two parameters to select an aerosol model from a suite of candidate models, and a third parameter to retrieve the corresponding aerosol optical depth. We use the parameter $\varepsilon_{ms}(\lambda, 865)$ defined as:

$$\varepsilon_{ms}(\lambda, 865) = [\rho_{path}(\lambda) - \rho_{ray}(\lambda)]/[\rho_{path}(865) - \rho_{ray}(865)]$$

to obtain initial estimates of the aerosol model. Here $\rho_{\textit{path}}(\lambda)$ is the atmospheric contribution (aerosols plus molecules) to the TOA reflectance, and $\rho_{\textit{ray}}(\lambda)$ is the contribution from molecules in the absence of aerosols. To finalize the aerosol model selection we employ a second parameter defined as:

$$\rho_{diff}(\lambda_i) = \frac{\rho_{tot}^{meas}(\lambda_j) - \rho_{tot}^{LUT}(\lambda_i)}{\rho_{tot}^{meas}(\lambda_j)} \times 100$$

where $\rho_{tot}^{LUT}(\lambda_j)$ is the TOA reflectance for wavelength λ_j , as inferred from the lookup tables (LUTs) based on the retrieved aerosol optical properties and chlorophyll concentration. The third parameter is defined as:

$$\gamma_{diff}(\lambda, 865) \equiv \gamma(\lambda) - \gamma(865)$$

where $\gamma(\lambda) = \rho_{path}(\lambda)/\rho_{ray}(\lambda)$.

The aquatic contribution to the TOA reflectance $t\rho_w(\lambda)$ is obtained simply by subtracting the inferred atmospheric contribution to the TOA reflectance $\rho_{tot}^{LUT}(\lambda)$ from the inferred total TOA reflectance $\rho_{tot}^{LUT}(\lambda)$: $t\rho_w(\lambda) = \rho_{tot}^{LUT}(\lambda) - \rho_{path}^{LUT}(\lambda)$. We use the ratio $R_w = t\rho_w(\lambda_1)/t\rho_w(\lambda_2)$ to retrieve the chlorophyll concentration, where $(\lambda_1 = 490 \text{ nm} \text{ and } \lambda_2 = 555 \text{ nm}$ (Stamnes *et al.*, 2003a).

In our iterative method for simultaneous retrieval of the chlorophyll concentration C and the aerosol optical properties (Stamnes et~al., 2003a), we use the parameters ε_{ms} (765,865) and γ_{diff} (765,865) to select the aerosol model, the aerosol optical depth τ (865), and the chlorophyll concentration the *first two times* through the algorithm iteration loop. Subsequently we use $\rho_{diff}(\lambda_i)$ to select that combination of aerosol model, aerosol optical depth, and chlorophyll concentration that gives the best match between inferred and measured reflectances for all available channels (Stamnes et~al., 2003a). Typically, 3-4 iterations are sufficient for convergence. Performance tests based on synthetic data show that our algorithm provides self-consistent and accurate retrievals, and applications to SeaWiFS data shows that it produces realistic results with quantifiable errors (Stamnes et~al., 2003a).

Extension of Algorithm to Coastal (Case 2) Waters

In order to adapt our algorithm for Case 1 waters to the more complex situation encountered in Case 2 coastal waters there are a number of issues that require attention. Some of these are: (i) How do we identify and deal with transition from Case 1 waters, where our current algorithm (hereafter called ALGO-C1 for short) applies, to Case 2 waters? (ii) How do we identify and deal with absorbing aerosols? (iii) How do we deal with sunglint and whitecaps? (iv) How do we deal with the bidirectional reflectance

distribution function (BRDF) caused by ocean waves? (v) How do we establish an error budget for retrievals in Case 2 waters?

To deal with absorbing aerosols we must establish an algorithm (called ALGO-C1A) for simultaneous retrieval of aerosol properties and chlorophyll concentration, when absorbing aerosols are present. In order to address the questions raised above we envision that the algorithm development will proceed in a number of steps approximately as follows:

- Step 0: For each image we start by establishing a Abs-no/Abs-yes mask that will allow us to distinguish between non-absorbing (Abs-no) and absorbing aerosols (Abs-yes), so that we can apply the proper algorithm AGLO-C1 or ALGO-C1A.
- Step 1: For a given scene, we use the ALGO-C1 algorithm (or ALGO-C1A as appropriate) to retrieve aerosol properties and chlorophyll concentration, C_{c_1} . If C_{c_1} 1.0 mg m⁻³, we classify this pixel as Case 1 water.
- Step 2: If $C_{c_1} > 1.0 \text{ mg} \cdot \text{m}^{-3}$, we use the ALGO-C2 algorithm to retrieve the properties of Case 2 water using the results from ALGO-C1 obtained in Step 1 as initial values of aerosol properties and $C_0 = C_{c_1}$. If the sediment concentration $S_{c_2} \varepsilon_1$, and $|C_{c_1} C_{c_2}| \le \varepsilon_2$, we set $C = C_{c_1}$, and classify this pixel as Case 1 water. Otherwise, we classify this pixel as Case 2 water, and S_{c_2} , C_{c_2} , and T_{c_2} are the retrieval results.
- Step 3: By repeating steps 1 and 2, we cover the entire scene.

Finally, we use the information obtained in steps 1 through 3 above as a *priori* information available to constrain the final retrieval, based on standard optimization techniques such as optimal estimation theory (Rodgers, 2000) or simulated annealing (Frette *et al.*, 1998; Frette *et al.*, 2001), of a complete set of aerosol properties and marine constituents in Case 2 waters that minimizes the difference between inferred and measured radiances. The output of this algorithm will be: (i) Marine Constituents in Case 2 waters (Chlorophyll, CDOM, Sediments), (ii) Aerosol Optical Properties (single scattering albedo and optical depth), and (iii) Estimates of Errors Associated with the Retrievals.

Summary

The main advantages of our new algorithm compared to current approaches appear to be fourfold: (i) many unnecessary assumptions invoked in current algorithms are avoided; (ii) our new algorithm is expected to perform more satisfactorily than current ones in situations with high relative humidity (Yan *et al.* 2002a; Stamnes *et al.*, 2003b); (iii) it is expected to be useful for addressing problems associated with absorbing aerosols as well as simultaneous retrieval of aerosol optical properties and marine constituents in Case 2 waters (Frette *et al.*, 1998; Frette *et al.*, 2001); (iv) it lends itself readily to the inclusion of more comprehensive IOP models, such as that presented by Stramski *et al.* (2001). Finally, a summary has been provided of the steps required to extend this algorithm to coastal waters.

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