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Westmoreland County, Virginia **Potomac River Shoreline**

Dune Evolution



Dune Evolution Westmoreland County, Virginia Potomac River Shoreline

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2006

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The views expressed herein are those of the authors and do not necessarily reflect the views of NOAA or any of its subagencies or DEQ.









TABLE OF CONTENTS

TABL	E OF CONT	ENTS i
LIST	OF FIGURES	5
LIST	OF TABLES	i
I.	INTRODUCA.GenerB.Chesa	CTION 1 ral Information 1 upeake Bay Dunes 1
II.	SHORE SET	FTING
	A.PhysiB.Hydro	cal Setting
III.	METHODSA.PhotoB.Rate of	6Rectification and Shoreline Digitizing6of Change Analysis6
IV.	RESULTS a	nd DISCUSSION of NEAR FUTURE TRENDS OF DUNE SITES
VI.	SUMMARY	
VII. Ackno	REFERENC wledgments .	ES
APPENDIX A APPENDIX B		Plates 3, 4, 8, 9, 10, 13, and 18 of Westmoreland's shoreline with historical aerial photography, digitized shorelines, and rates of shoreline change. Tables of specific dune site information.

LIST OF FIGURES

- Figure 1. Location of Westmoreland County within
- Figure 2. Location of localities in the Dune Act with
- Figure 3. Index of shoreline plates
- Figure 4. Geological map of Westmoreland County
- Figure 5. Variability of dune and beach profiles with Figure 6. Typical profile of a Chesapeake Bay dune
- Figure 7. Ground photos of select Westmoreland dur

Table 1. Summary wind conditions at Quantico Mart Table 2. Summary shoreline rates of change and their

Cover Photo: Photograph of site WM18, the Murphy headland breakwater system on the Potomac River. Photo taken by Shoreline Studies Program on 25 September 2002.

Bay estuarine system
al., 1989)
d County
d County.

LIST OF TABLES

rine Corps Base from	1973-2001	5
ir standard deviation		1

I. INTRODUCTION

A. General Information

Shoreline evolution is the change in shore position through time. In fact, it is the material resistance of the coastal geologic underpinnings against the impinging hydrodynamic (and aerodynamic) forces. Along the shores of Chesapeake Bay, it is a process-response system. The processes at work include winds, waves, tides and currents, which shape and modify coastlines by eroding, transporting and depositing sediments. The shore line is commonly plotted and measured to provide a rate of change but it is as important to understand the geomorphic patterns of change. Shore analysis provides the basis to know how a particular coast has changed through time and how it might proceed in the future.

The purpose of this report is to document how the dunes on Potomac River shores of Westmoreland (Figure 1) have evolved since 1937. Aerial imagery was taken for most of the Bay region beginning that year, and it is this imagery that allows one to assess the geomorphic nature of shore change. Aerial imagery shows how the coast has changed, how beaches, dunes, bars, and spits have grown or decayed, how barriers have breached, how inlets have changed course, and how one shore type has displaced another or has not changed at all. Shore change is a natural process but, quite often, the impacts of man through shore hardening or inlet stabilization come to dominate a given shore reach. Most of the change in shore positions where dunes occur will be quantified in this report. Others, particularly very irregular coasts, around inlets, and other complicated areas will be subject to interpretation.

B. Chesapeake Bay Dunes

The primary reason for developing this Dune Evolution report is to be able to determine how dunes and beaches along the coast of Westmoreland have and will evolve through time. The premise is that, in order to determine future trends of these important shore features, one must understand how they got to their present state. Beaches and dunes are protected by the Coastal Primary Sand Dune Protection Act of 1980 (Act)¹. Research by Hardaway *et al.* (2001) located, classified and enumerated jurisdictional dunes and dune fields within the eight localities listed in the Act. These include the counties of Accomack, Lancaster, Mathews, Northampton and Northumberland and the cities of Hampton, Norfolk and Virginia Beach (Figure 2). Only Chesapeake Bay and river sites were considered in that study.

In 2002, Hardaway *et al.* characterized dunes in several non-jurisdictional localities including Westmoreland. That report detailed the location and nature of the jurisdictional primary dunes along the River shore of Westmoreland, and those results appear in Appendix B. For this study, the positions of the dune sites are presented using the latest imagery in order to see how the sites sit in the context of past shoreline positions. The dune location information has not been field verified since the original visits in 2001. This information is not intended to be used for jurisdictional determinations regarding dunes.

¹The General Assembly of Virginia enacted the Coastal Primary Sand Dune Protection Act (the Dune Act) in 1980. The Dune Act was originally codified in § 62.1-13.21 to -13.28. The Dune Act is now recodified as Coastal Primary Sand Dunes and Beaches in § 28.2-1400 to -1420.

II. SHORE SETTING

A. Physical Setting

The Potomac River of Westmoreland County extends from the county line with Northumberland at the Yeocomico River upriver to Rosier Creek and the border with King George County (Figure 3). This includes about 40 miles of tidal shoreline along the Potomac River. Historic shore erosion rates vary from 0 ft/yr to over 4 ft/yr along the Potomac River shoreline (Byrne and Anderson, 1978).

The coastal geomorphology of the County is a function of the underlying geology and the hydrodynamic forces operating across the land/water interface, the shoreline. The Potomac River coast of Westmoreland is undifferentiated Lynnhaven and Poquoson Members in the lower section of the county to Nomini Bay. Along a short section of the Potomac river, the older Tertiary Chesapeake Group outcrops with high banks (Figure 4). The Quaternary Sedgefield Member occurs farther upriver. The younger Quaternary geology is generally lower in elevation than the older Tertiary formations. The Atlantic Ocean has come and gone numerous times over the Virginia coastal plain over the past million years or so creating these various geologic formations. The effect has been to rework older deposits into beach and lagoonal deposits at time of the transgressions. The last low stand found the ocean coast about 60 miles to the east when sea level about 300 feet lower than today and the coastal plain was broad and low. The current estuarine system was a meandering series of rivers working their way to the coast. About 15,000 years ago, sea level began to rise and the coastal plain watersheds began to flood. Shorelines began to recede. The slow rise in sea level is one of two primary long-term processes which cause the shoreline to recede; the other is wave action, particularly during storms. As shorelines recede or erode, the bank material provides the sands for the offshore bars, beaches and dunes. Parts of Westmoreland's littoral system is sand rich from erosion over time of the sandy, sometimes high, upland banks and the nearshore substrate.

Sea level is continuing to rise in Chesapeake Bay. Tide data collected at Sewells Point in Norfolk show that sea level has risen 4.42 mm/yr (0.17 inches/yr) or 1.45 ft/century (http://www.co-ops.nos.noaa.gov/). Lewisetta on the Potomac River in neighboring Northumberland County rose 4.85 mm/yr or 1.59 ft/century. Increased water levels directly effect the reach of storms and their impact on shorelines. Anecdotal evidence of storm surge during Hurricane Isabel, which impacted North Carolina and Virginia on September 18, 2003, put it on par with the storm surge from the "storm of the century" which impacted the lower Chesapeake Bay in August 1933. Boon (2003) showed that even though the tides during the storms were very similar, the difference being only 4 cm (~0.5 in), the amount of surge was different. The 1933 storm produced a storm surge that was greater than Isabel's by slightly more than a foot. However, analysis of the mean water levels for the months of both August 1933 and September 2003 showed that sea level has risen by 41 cm (1.35 ft) at Hampton Roads in the seventy years between these two storms (Boon, 2003). This is the approximate time span between our earliest aerial imagery (1937) and our most recent (2002), which means the impact of sea level rise to shore change is significant. The beaches, dunes, and nearshore sand bars try to keep pace with the rising sea levels.







Figure 2. Location of localities in the Dune Act with jurisdictional and non-jurisdictional localities noted.



Figure 3. Westmoreland plate index.

Plates with dune sites and shown in Report
Plates not shown in Report



Qs	Holocene Sand - Pale gray to light-yellowish gray, fine to coarse, poorly sorted to well sorted, shelly in part; contains angular to rounded fragments and whole valves of mollusks. Comprises deposits of coast narrow beach-dune ridges bordering brackish-water marshes of Chesapeake Bay. As much as 40 ft in thickness.
Qtip	Lynnhaven and Poquoson Members, undifferentiated.
Qsh	Shirley Formation (middle Pleistocene) - Light-to dark-gray and brown sand, gravel, silt, clay, and peat. Constitutes surficial deposits of riverine terraces and relict baymouth barriers and bay-floor plains (alt. 35-45 ft) inset below the Chuckatuck Formation (Johnson and Peebles, 1984). Upper part of unit is truncated on the east by the Suffolk and Harpersville scarps; locally, lower part extends east of scarps. Fluvial-estuarine facies pebble to boulder sand overlain by (2) fine to coarse sand interbedded with peat and clayey silt rich in organic material, including in situ tree stumps and leaves and seeds of cypress, oak, and hickory, which medium- to thick-bedded, clayey and sandy silt and silty clay. Marginal-marine facies in lower James River and lowermost Rappahannock River areas is silty fine sand and sandy silt containing Crassostrea Mercenaria, and other mollusks. Astrangia from lower Rappahannock River area has yielded a uranium-series age of 184,000 +/- 20,000 years B.P. (Mixon and other, 1982). Thickness is 0-80 ft.
QTw	Windsor Formation (lower Pleistocene or upper Pliocene) - Gray and yellow to reddish-brown sand, gravel, silt, and clay. Constitutes surficial deposits if extensive plain (alt. 85-95 ft) seaward of Su fluvial-estuarine terrace west of scarp. Fining-upward sequence beneath plain consists of basal pebbly sand grading upward into crossbedded, quartzose sand and massive, clayey silt and silty clay; lo sequence were deposited, respectively, in shallow-marine or open-bay and restricted-bay or lagoonal environments. In terraces west of Surry scarp, fluvial-estuarine deposit comprise Muddy, coarse, trans and gravel grading upward to sandy silt and clay. Unit is 0-40 ft thick.
Qts	Sedgefield Member - Pebbly to bouldery, clayey sand and fine to medium, shelly sand grading upward to sandy and clayey silt; locally, channel fill at base of unit includes as much as 50 ft of fine to coarse, crossbedded sand and c peat containing in situ tree stumps. Sandy bay facies commonly contains Crassostrea biostromes, Mercenaria, Anadara, Polynices, Ensis, and other mollusks. Specimes of the coral Astrangia have yielded estim ages averaging 71,000 +/- 7,000 yrs B.P. (Mixon and others, 1982). Unit constitutes surficial deposit of river- and coast-parallel plains (alt. 20-30 ft) bounded on landward side by Suffolk and Harpersville scarps. The
Tc Tm	Chesapeake Group (upper Pliocene to lower Miocene) - Fine to coarse, quartzose sand, silt, and clay; variably shelly and diatomaceous, deposited mainly in shallow, inner- and middle-shelf waters in studies of foraminiferal, nannofossil, diatom, and molluscan assemblages in Virginia and adjacent states (Andrews, 1988; Gibson, 1983; Gibson and others, 1980; Poag, 1989; Ward and Blackwelde Krafft, 1984), Includes the following formations (see also sheet 2, figure 1), from youngest to oldest; Chowan River Formation (upper Pliocene), Yorktown Formation (lower upper and lower Pliocene), (upper Miocene), St. Mary's Formation (upper and Moorings unit of Oaks and Coch (1973) (upper Pliocene) - White, light-gray, and grayish-yellow quartzose sand and gray to grayish-brown clayed silt and silty clay. Constitutes discontinuous linear scrap; depositional. Surfaces range in altitude from 130 ft along slightly higher, ridge-like topography at scarp to about 110 ft west of scarp. Eastern facies of unit is unfossiliferous, massive to cross-law fine sand.
Tb ¹ Tb ²	 Bacons Castle Formation (upper Pliocene) - Gray, yellowish-orange, and reddish-brown sand, gravel, silt, and clay; Constitutes surficial deposits of high plain extending from Richmond, VA., Eas subdivided into two members: Tb¹ and Tb². Tb² predominantly thin-bedded and laminated clayey silt and silty fine sand. TB² is characterized by flaser, wavy, and lenticular bedding and rare to <i>Ophiomorpha nodosa</i>. Unit is 0-70 ft thick. Pliocene sand and Gravel - Interbedded yellowish-orange to reddish-brown gravelly sand, sandy gravel, and fine to coarse sand, poorly to well-sorted, crossbedded in part, includes less amounts o medium beds. Commonly caps drainage divides (alt. 250-170ft) in northwestern part of Coastal Plain. Lower part of unit, showing flaser and lenticular bedding and contain rare to abundant <i>Ophiomor</i> deposition in marginal-marine environment and is in part, a nearshore equivalent of the more downslip, marine facies of the Yorktown Formation. The more fluvial-deltaic sediments that prograded ead during a regressive phase of the Yorktown. Thickness is 0-50ft.

Figure 4. Geologic map of Westmoreland County (from Mixon et al., 1989).

tal barrier islands and

 depositional surfaces of s comprises (1) a lower
 grades upward to (3)
 a virginica, Mulinia, Noetia,

Surry scarp and coeval, lower and upper parts of trough-crossbedded sand

clayey silt and mated uranium-series Fhickness is 0-50 ft.

ers. Ages of units based der, 1980; Ward and), Eastover Formation

r body along and just west of Surry aminated, moderately well-sorted,

stward to Surry Scrap. Unit is common clay-lined burrows including

of clay and silt in thin to orpha nodosa, represents astward across the shelf

WIND DIRECTION										
Wind Speed (mph)	Mid Range (mph)	North	North east	East	South east	South	South west	West	North west	Total
< 5	3	5703* 3.21 ⁺	3330 1.87	3868 2.18	4792 2.70	12257 6.90	4291 2.42	7070 3.98	15437 8.69	56748 31.95
5-11	8	17454 9.82	10087 5.68	6504 3.66	8117 4.57	22593 12.72	8515 4.79	13391 7.54	18453 10.39	105114 59.17
11-21	16	3698 2.08	1460 0.82	386 0.22	517 0.29	2030 1.14	1156 0.65	1129 0.64	4601 2.59	14977 8.43
21-31	26	165 0.09	64 0.04	34 0.02	21 0.01	60 0.03	64 0.04	102 0.06	274 0.15	784 0.44
31-41	36	7 0	1 0	2 0	0 0	1 0	1 0	7 0	7 0	26 0.01
41-50	46	0 0	0 0	0 0	0 0	1 0	0 0	0 0	0 0	1 0
>50		1 0	3 0	3 0	3 0	4 0	0 0	7 0	5 0	26 0.01
Total		27028 15.20	14945 8.41	10797 6.08	13450 7.57	36946 20.79	14027 7.9	21706 12.22	38777 21.82	177676 100.00

B. Hydrodynamic Setting

Mean tide range along the upper Potomac River coast of Westmoreland is about 1.2 ft (1983-2001 Tidal Epoch at Lewisetta). Spring tide range is 1.5 ft. The wind/wave climate impacting the Westmoreland shore is defined by fetch exposures to the northwest, north, northeast, and east along Potomac River. Wind data from Quantico Marine Corps Base upriver reflect the frequency and speeds of wind occurrences from 1973 to 2001 (Table 1) which characterize the locally-generated Bay waves.

Northeasters are particularly significant in terms of the impacts of storm surge and waves on beach and dune erosion. Hurricanes, depending on their proximity and path can also have an impact to the Virginia coast. On September 18, 2003, Hurricane Isabel passed through the Virginia coastal plain. The main damaging winds began from the north and shifted to the east then south. Beach erosion and dune scarping were significant but areas with wide beaches offered more protection to the adjacent dunes.

*Number of occurrences ⁺Percent

Table 1. Summary wind conditions at Quantico Marine Corps Base from 1973-2001.

III. METHODS

A. Photo Rectification and Shoreline Digitizing

Recent and historic aerial photography was used to estimate, observe, and analyze past shoreline positions and trends involving shore evolution for Westmoreland. Some of the photographs were available in fully geographically referenced (georeferenced) digital form, but most were scanned and orthorectified for this project.

Aerial photos from VIMS Shoreline Studies as well as from United States Geological Survey (USGS) archives were acquired. The years used for the shoreline change analysis included 1937, 1969, 1987, 1994, and 2002. Color aerials were obtained for 1994 and 2002. The 1994 imagery was processed and mosaicked by USGS, while the imagery from 2002 was acquired from the Virginia Base Mapping Program (VBMP). The aerial photography for the remaining years were mosaicked by the VIMS Shoreline Study Program.

The images were scanned as tiffs at 600 dpi and converted to ERDAS IMAGINE (.img) format. They were orthorectified to a reference mosaic, the 1994 Digital Orthophoto Quarterquadrangles (DOQQ) from USGS. The original DOQQs were in MrSid format but were converted into .img format as well. ERDAS Orthobase image processing software was used to orthographically correct the individual flightlines using a bundle block solution. Camera lens calibration data was matched to the image location of fiducial points to define the interior camera model. Control points from 1994 USGS DOQQ images provide the exterior control, which is enhanced by a large number of image-matching tie points produced automatically by the software. A minimum of four ground control points were used per image, allowing two points per overlap area. The exterior and interior models were combined with a 30-meter resolution digital elevation model (DEM) from the USGS National Elevation Dataset (NED) to produce an orthophoto for each aerial photograph. The orthophotographs that cover each USGS 7.5 minute quadrangle area were adjusted to approximately uniform brightness and contrast and were mosaicked together using the ERDAS Imagine mosaic tool to produce a one-meter resolution mosaic also in an .img format.

To maintain an accurate match with the reference images, it was necessary to distribute the control points evenly. This can be challenging in areas with little development. Good examples of control points are permanent features such as manmade features and stable natural landmarks. The maximum root mean square (RMS) error allowed is 3 for each block.

Once the aerial photos were orthorectified and mosaicked, the shorelines were digitized in ArcMap with the mosaics in the background to help delineate and locate the shoreline. For Westmoreland's coast, an approximation to mean low water (MLW) was digitized. This is approximately the edge of the marsh or the "toe" of the beach. In areas where the shoreline was not clearly delineated on the aerial photography, the location was estimated based on the experience of the digitizer. Digitizing the shoreline brings in, perhaps, the greatest amount of potential error because of the problems of image clarity and definition of shore features. A series of Westmoreland dune site profiles are displayed in Figure 5 which shows beach/dune variability. Figure 6 shows the relationship of MHW, MLW and beach/dune system components.

B. Rate of Change Analysis

A custom Arcview extension called "shoreline" was used to analyze shoreline rate of change. A straight, approximately shore parallel baseline is drawn landward of the shoreline. The extension creates equally-spaced transects along the baseline and calculates distance from the baseline at that location to each year's shoreline. The output from the extension are perpendicular transects of a length and interval specified by the user. The extension provides the transect number, the distance from beginning baseline to each transect, and the distance from the baseline to each digitized shoreline in an attribute table. The attribute table is exported to a spreadsheet, and the distances of the digitized shoreline from the baseline are used to determine the rates of change. The rates of change are summarized as mean or average rates and standard deviations for each Plate.

It is very important to note that this extension is only useful on relatively straight shorelines. In areas that have unique shoreline morphology, such as creek mouths and spits, the data collected by this extension may not provide an accurate representation of true shoreline change. The shore change data was manually checked for accuracy. However, where the shoreline and baseline are not parallel, the rates may not give a true indication of the rate of shoreline change.



Figure 5. Variability of dune and beach profiles in Westmoreland County.



Figure 6. Typical profile of a Chesapeake Bay dune system (from Hardaway et al., 2001).

IV. RESULTS and DISCUSSION of NEAR FUTURE TRENDS OF DUNE SITES

The Plates referenced in the following sections are in Appendix A. Dune locations are shown on all photo dates for reference only. Dune sites and lengths are positioned accurately on the 2002 photo. Because of changes in coastal morphology, the actual dune site might not have existed earlier. Site information tables are in Appendix B. More detailed information about Chesapeake Bay dunes and individual dune sites in Westmoreland can be found in Hardaway *et al.* (2001) and Hardaway *et al.* (2002). Only those Plates that have dune sites are shown in Appendix A.

The shoreline trends are delineated based on past performance. Ongoing shore development, shore stabilization and/or beach fill, and storms will have local impacts on the near term. "Near Future" is quite subjective and only implies a reasonable expectation for a given shore reach to continue on its historic course for the next 10 to 20 years. In addition, the basis for the predictions are the shorelines digitized on geo-rectified aerial photography which have an error associated with them (see Methods, Section III). Each site's long-term and recent stability as well as a near future prediction are shown in a table in Appendix B. This data is intended as a resource for coastal zone managers and homeowners; it is not intended for use in determining legal jurisdictional limits.

Westmoreland County's Potomac River coast extends from the Yeocomico River northward to Rosier Creek just upriver of the Town of Colonial Beach. Twenty-two plates subdivide the shore along the Potomac River but only 7 plates have dune sites. These are plates 3, 4, 8, 9, 10, 13, and 18; only these plates shall be discussed in this section. Plate 3 has one dune site, WM5, that is located on the upriver side of the north jetty at Bonum Creek. It has evolved primarily due to the jetty trapping littoral sands. Shore change along Plate 3 has been quite variable over time with a net rate of -1.9 ft/yr. Site WM5 appears to be stable for the near term.

Site WM6 is the only dune site on Plate 4 and it lies at the mouth of Jackson Creek. The site was a spit feature in 1937 and appears to have gained backshore vegetation by 1969. Aerial imagery in 1987 shows a dredged channel and dredge material deposition on the downriver side of Jackson Creek. Site WM6 further evolved on the river side of that deposition. The short length of shoreline measured on Plate 4 includes the shoreline on either side of Jackson Creek and shows a net recession rate of -2.2 ft/yr. Since WM6 is backed by sandy dredge material against a maintained channel, it should be relatively stable.

Plate 8 goes from Glebe Point to Cabin Point along Lower Machodoc Creek and has three dune sites, WM14 and WM15/16. Site WM14 is an erosional remnant of a longer, more extensive beach and spit as seen in 1937 imagery. The Glebe Point spit has been reduced to almost no shore at all in 2002. WM14 marks the mainland attachment position of the spit. Site WM14 has had a history of erosion and that trend will most likely continue until it disappears.

The inlet to Cabin Point Creek was moved upriver by dredging a new channel in the late 1980s and can be seen in 1994 imagery. Site WM15/16 evolved in the groinfield upriver that was installed when the channel was created. The shoreline has become relatively stable due to the groins; however the source of sand to the reach has been cut off by extensive bulkheading upriver. The site may experience reduced sand supply with consequent reduction in beach width and dune erosion.

Sites WM18A and WM18B, on Plate 9, exists because of the offshore headland breakwater installed in 1995. The site evolved on the attached tombolo. The headland breakwaters allow the adjacent banks to erode into an equilibrium planform. This site may grow with time as more sand enters the embayments from the eroding sandy

banks as well as from onshore deposition. Overall, the net long-term erosion on the Plate 9 coast is -1.8 ft/yr.

Plate 10 had one dune site as well, WM20. It is part of a long, stable curvilinear embayment formed, in part, by a long jetty at the mouth of Nomini Creek. The net shoreline change rate is -0.6 but the embayed coast from 4,000 to 11,000 is closer to +0.5 ft/yr. This is a very stable coast, and WM20 should remain in existence as long as the jetty is maintained.

Plate 13 has one dune site, WM24, that has evolved as the south spit on Hollis Marsh has grown and decayed over time. Hollis Marsh island has diminished in size, in length, as well as breadth over time. Site WM24 was in its present position by 1987, increased in length, then retreated slightly by 2002. It is stable for now, but further breakup of Holly Marsh will impact its fate. Shore change patterns show erosion of the Hollis Marsh Island coast with accretion then erosion of the southern spit. The net rate of change for the island is -5.9 ft/yr.

Plate 18 has dune site WM28 which lies in a slightly curved, embayed coast at George Washington's Birthplace National Monument. A narrow beach and possibly a dune can be seen in 1937 imagery. A small un-named creek inlet also can be seen which has an associated ebb shoal that has helped maintain a beach and dune on its landward flanks. In fact, that area has a net (1937-2002) accretion rate of +0.3 ft/yr. However, the shoal and beaches have been reduced over time, and this trend appears to be continuing.





Figure 7. Ground photos of select Westmoreland dune sites.

VI. SUMMARY

Shoreline change rates are based on aerial imagery taken at a particular point in time. For those plates shown in the report, the rate of change was calculated every 500 ft. The mean or average rate for each plate is shown in Table 2 for five time periods with the long-term rate determined between 1937 and 2002. The total average and standard deviation (Std Dev) for the entire data set of individual rates is also given. The standard deviation shows the relative spread of values about the mean or average. Larger standard deviation values relative to the mean indicates a wider scatter of erosion rates about the mean while lower standard deviation values indicates erosion rates are concentrated near the mean (*i.e.* all the rates calculated for the entire plate were similar).

The standard deviations are close to the average rate of change indicating that the shore change rates were relatively consistent for that time period. Overall rates of change were relatively small with the exception of Plate 13 which depicts Hollis Marsh Island separating Currioman Bay from the Potomac River. This section of shore has been highly erosive since 1937, and the negative shore change rate has been increasing ever since. Plates 8, 10,and 18 had the lowest overall rates of change. However, for Plates 8 and 10, the standard deviation is much larger than the average rate of change indicating that the overall rate is probably not indicative of the change which occurred on this section of shore. When short time frames are used to determine rates of shoreline change, shore alterations may seem amplified. The rates based on short-time frames can modify the overall net rates of change.

The shore change patterns shown in this report along with the aerial imagery will indicate how the coast will evolve based on past trends and can be used to provide the basis for appropriate shoreline management plans and strategies. Dunes and beaches are a valuable resource that should be either maintained, enhanced or created in order to abate shoreline erosion and provide sandy habitat.

	Plate 3		Plate 4		Plate 8		Plate 9		Plate 10		Plate 13		Plate 18	
Imagery Dates	Rate of Change (ft/yr)	Std. Dev.												
1937-1969	-1.4	1.1	-1.3	1.1	-0.4	2.4	-1.8	0.8	-1.1	2.9	-4.9	5.4	0.3	1.5
1969-1987	-2.8	1.9	-3.4	2.5	1.1	3.5	-1.0	2.5	0.1	1.6	-5.4	9.2	-0.4	0.9
1987-1994	-2.6	2.5	-2.5	2.0	-3.0	3.8	-1.4	2.3	0.4	1.3	-6.4	8.1	-2.8	1.4
1994-2002	-1.3	1.9	-2.1	2.2	-2.7	3.6	-3.6	3.6	-0.9	1.4	-10.4	14.5	-4.1	1.5
1937-2002	-1.9	1.0	-2.2	1.0	-0.7	1.8	-1.8	1.2	-0.6	1.5	-5.9	6.4	-0.6	0.9

Table 2. Summary shoreline rates of change and their standard deviation for Westmoreland County.

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APPENDIX A

For each Plate shown on Figure 3, Appendix A contains orthorectified aerial photography flown in 1937, 1969, 1987, 1994, and 2002. Also shown are the digitized shorelines, identified dune sites, and an arbitrarily created baseline.
A plot shows only the relative locations of the shorelines while another one depicts the rate of shore change between dates.
A summary of the average Plate rate of change in ft/yr as well as the standard deviation for each rate is also shown.

This data is intended as a resource for coastal zone managers and homeowners; it is not intended for use in determining legal jurisdictional limits.

Plate 3	Plate 4	Plate 8
Plate 9	Plate 10	
Plate 13	Plate 18	











A5





A7











Westmoreland County Plate 8





00 0

1,000 Feet



























2002

Westmoreland County Plate 13













Westmoreland County Plate 18





APPENDIX B

The data shown in the following tables were primarily collected as part of the Chesapeake Bay Dune Systems: Monitoring Year One report and presented in Hardaway *et al.* (2002). Individual site characteristics may now be different due to natural or man-induced shoreline change.

An additional table presents the results of this analysis and describes each dune site's relative long-term, recent, and near-future predicted stability. This data results from the position of the digitized shorelines which have an error associated with them (see Methods, Section III).

Since much of the dune data were collected several years ago and the beach and dune systems may have changed, this report is intended as a resource for coastal zone managers and homeowners; it is not intended for use in determining legal jurisdictional limits.

These data were collected as part of the Chesapeake Bay Dune Systems: Monitoring Year One (Hardaway et al., 2002). Site characteristics may now be different due to natural or man-influenced shoreline change.

	Location [^]			Dune	Primary	Secondary	*Public
Dune				Shore	Dune	Dune	Ownership?
Site No.	Easting	Northing		Length	Site?	Site?	
	(feet)	(feet)	Date Visted	(feet)			
5	2551770	647818	16-Aug-2001	170	Yes		
6	2547691	651065	16-Aug-2001	320	Yes	Yes	
14	2531440	660758	27-Sep-2001	400	Yes		
15/16	2528489	663735	29-Nov-2001	800	Yes		
18a	2521524	669004	29-Nov-2001	90	Yes		
18b	2521524	669004	29-Nov-2001	90	Yes		
22	2512881	662546	29-Nov-2001	1200	Yes		
24	2506253	665760	29-Nov-2001	1000	Yes		
28	2450953	683746	27-Sep-2001	200	Yes		Yes
Total				4270			

Identified dune sites in Westmoreland County as of date of site visit.

*Public ownership includes governmental entities including local, state, and federal;

otherwise ownership is by the private individual.

^Location is in Virginia State Plane South, NAD 1927

Dune site measurements in Westmoreland County as of date of site visit.

				Dune Site Measurements							
		Dune		Primary Dune		Secondary Dune					
		Shore	Crest	Distance fro	om Crest			Dist	ance From		
		Length	Elev	Landward	To MLW	Juris-	Crest	Primary Crest	2ndCrest	2nd Crest seaward	
	Site			to Back Base		diction	Elev	Landward to Profile end	landward	to 1st back base	
	No.	(Feet)	(ft MLW)	(Feet)	(Feet)		(ft MLW)	(Feet)	(Feet)	(Feet)	
WM	5	170	4.27	16	42.7						
WM	6	320	5.06	28	44.1	Y	4.67	90	23	34	
WM	14	400	5.94	16	42.9						
WM	15/16	800	5.83	35	64.9						
WM	18a	90	3.54	7	83.6						
WM	18b	90	5.78	29	66.9						
WM	22	1200	5.51	52	60.5						
WM	24	1000	3.69	51							
WM	28	200	5.02	4.5	56.8						

These data were collected as part of the Chesapeake Bay Dune Systems: Monitoring Year One (Hardaway *et al.*, 2002). Site characteristics may now be different due to natural or man-influenced shoreline change.

Dune site parameters in Westmoreland County as of date of site visit.

				Dune Site Parameters						
			Fetch	Shoreline	Near	shore	Morphologic	Relative	Underlying	Structure
			Exposure	Direction	Grad	dient	Setting	Stability	Substrate	or Fill
	Site	Туре		of Face						
	No.		А	В	(C	D	E	F	G
WM	5	Man Inf	Riverine	North	Steep	No Bars	Isolated, Linear	Stable	Marsh/CB	Jetty
WM	6	Man Inf	Riverine	Northeast	Steep	No Bars	Isolated, Linear	Stable	Marsh/CB	Groin
WM	14	Man Inf	Riverine	Northeast	Medium	No Bars	Creek Mouth Barrier/ Spit	Erosional	Marsh/CB	Groin
WM	15/16	Man Inf	Riverine	Northeast	Medium	No Bars	Creek Mouth Barrier/ Spit	Stable	Marsh/CB	Groin, Jetty
WM	18a	Man Inf	Riverine	East	Steep	No Bars	Isolated, Linear	Stable	Upland	BW
WM	18b	Man Inf	Riverine	Northwest	Steep	No Bars	Isolated, Linear	Stable	Upland	BW
WM	22	Man Inf	Riverine	Northwest	Medium	No Bars	Creek Mouth Barrier/ Spit	Stable	Marsh/CB	Jetty
WM	24	Man Inf	Riverine	East	Medium	No Bars	Creek Mouth Barrier/ Spit	Erosional	Marsh/CB	
WM	28	Natural	Riverine	North	Steep	No Bars	Isolated, Linear	Stable	Upland	

Long-term, recent stability and future predictions of shore erosion and accretion rates for dune sites in Westmoreland County.

Site	No.
WM	5
WM	6
WM	14
WM	15/16
WM	18
WM	20
WM	24
WM	28

Long-Term	Recent	Near
Stability	Stability	Future
1937-2002	1994-2002	Prediction
Accretionary	Stable	Stable
Stable	Stable	Stable
Erosional	Erosional	Erosional
Erosional	Stable	Erosional
Erosional	Stable	Stable
Accretionary	Erosional	Stable
Accretionary	Accretionary	Stable*
Stable	Erosional	Erosional

*Will eventually erode