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1975

Pleistocene geology of the Grand Forks-Bemidji area, northwestern Minnesota

Kenneth L. Harris *University of North Dakota*

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PLEISTOCENE GEOLOGY OF THE GRAND FORKS-BEMIDJI AREA.

NORTHWESTERN MINNESOTA

by

Kenneth L. Harris

Bachelor of Science, North Dakota State University, 1969 Master of Science, University of North Dakota, 1973

^ADissertation

Submitted to the Graduate Faculty

of the

University of North Dakota

in partial fulfillment of the requirements

for the degree of

Doctor of Philosophy

Grand Forks, North Dakota

May 1975

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ACKNOWLEDGMENTS

The members of my committee are gratefully acknowledged for their constructive criticism and discussion during the course of this study. In particular the critical evaluations of Lee Clayton, Stephen R. Moran, Walter L. Moore, and John A. Brophy were appreciated.

Several people conducting related studies in the region offered helpful suggestions and criticism. They include Mark M. Fenton of the Geology Department of the University of Western Ontario, Canada; Donald K. Sackreiter and Howard C. Hobbs of the Geology Department of the University of North Dakota; and B. Michael Arndt of the North Dakota Geological Survey. Stephen R. Moran, Donald K. Sackreiter and B. Michael Arndt are also gratefully acknowledged for offering aid and assistance during the course of the subsurface drilling program.

Edwin A. Noble and the North Dakota Geological Survey are gratefully acknowledged for making available laboratory facilities, miscellaneous supplies, and a truck mounted power-auger necessary for the completion of this study.

Matt Walton and the Minnesota Geological Survey are gratefully acknowledged for their assistance in the subsurface drilling program and other phases of this study.

My wife Phyllis offered the most aid and comfort during the course of this study.

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ABSTRACT

This is a report of a reconnaissance study of the Quaternary geology of an area in northwestern Minnesota. Surface geology was studied in an area extending from Grand Forks, North Dakota, in the northwest to Bemidji, Minnesota, in the southeast, an area of 10 723 square kilometres (4140 square miles). Near-surface stratigraphy was studied in the area of surface study and an adjacent area of equal size to the south.

Surface materials mapped range in age from Pleistocene to Holocene and include glacial, glaciofluvial, lacustrine, bog, eolian, and alluvial sediments.

Sixteen power-auger test holes, two measured sections along the Red Lake River, two composite sections from earlier drilling programs and surface exposures were used to characterize seven nearsurface lithostratigraphic units in the area. From oldest to youngest they are unnamed unit 1, unnamed unit 2, the Marcoux Formation, and the St. Hilaire Formation. These units are largely glacial sediment and pre-Wisconsinan or Early Wisconsinan in age. The Red Lake Falls Formation and the Huot Formation are largely glacial sediment and Wisconsinan and latest Wisconsinan in age. The Sherack Formation is largely lacustrine sediment and is latest Wisconsinan and Holocene in age. The lithostratigraphic units present are differentiated by means of their texture and coarse-sand lithology.

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During the Pleistocene an unknown number of glaciations occurred before the deposition of the glacial sediments of the oldest lithostratigraphic unit observed, unnamed unit 1. Three pre-Wisconsinan or Early Wisconsinan glaciers deposited the glacial sediments of unnamed unit 1, unnamed unit 2, and the Marcoux Formation over the entire area. Wisconsinan glaciers advanced into the area, flowed around the Itasca Highland, and retreated twice, depositing the glacial sediment of the St. Hilaire and Red Lake Falls Formations. The late Wisconsinan and Holocene lake sediments of the Argusville, Wylie, and Sherack Formations were deposited in Lake Agassiz during the retreat of the glacier that deposited the Red Lake Falls Formation and the advance and retreat of the glacier that deposited the Huot Formation. Lake Agassiz drained at about 9500 BP.

INTRODUCTION

Purpose of Study

This report is a summary of a reconnaissance of the Quaternary geology of an area in northwestern Minnesota. The surface geology of the study area was mapped, and the near-surface stratigraphy was studied. The information and interpretations from this study and from other studies in adjacent parts of Minnesota, North Dakota, and Manitoba provide insight into the Pleistocene history of the Upper Midwest.

Area of Study

Two different areas are considered here: a surface study area and a subsurface study area. The Grand Forks-Bemidji area, the area of surface study, is located in northwestern Minnesota. It is bounded on the west by the Red River of the North and extends east to 94° 30' west longitude. The area is bounded on the north by 48° north latitude and extends to the south to 47° 30' north latitude (Figure 1). The area of surface study contains about 115 townships or about 10 723 square kilometres (4140 square miles). Parts of Polk, Pennington, Clearwater, and Beltrami Counties and all of Red Lake County are included in the study area.

The area of subsurface study in this report consists of the Grand Forks-Bemidji area and an adjacent area of equal size to the south. The subsurface study area is bounded on the west by the Red

Fig. 1. Location of the Grand Forks-Bemidji Area (crosshatched) within the north half of the subsurface study area.

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River of the North and extends east to 94° 30' west longitude. The subsurface study area is bounded on the north by 48° north latitutde and on the south by 47° north latitude (Figure 1).

The Grand Forks-Bemidji area contains a variety of geomorphic features. The Lake Agassiz plain and shoreline complex dominate the western and northern parts of the area. Collapsed glacial topography and outwash features are present in the central and eastern parts of the area.

The study area spans the boundaries between the prairie in the west, deciduous forest in the central part, and conifer-hardwood forest in the east.

Previous Work

Two early geologists who studied at least parts of the study area were Warren Upham and Frank Leverett. Warren Upham did reconnaissance geologic mapping of Minnesota from 1872 to 1895. He and H. H. Winchell, the first Minnesota State Geologist, published papers that dealt mainly with descriptions and interpretations of local landforms (Wright, 1972). Upham's major contribution to the understanding of the Pleistocene history of the region was his study of glacial Lake Agassiz (Upham, 1896).

Frank Leverett, a United States Geological Survey geologist, worked in Minnesota during the period from 1906 to 1912. Leverett's surficial maps of Minnesota are still the only surficial maps of the entire state (Leverett, 1915; Leverett and Sardeson, 1917; Leverett and Sardeson, 1919). Leverett summarized his work in a United States

Geological Survey professional paper, Quaternary geology of Minnesota and parts of adjacent states (Leverett, 1932).

A 1933 publication by C. C. Nikiforoff, a United States Department of Agriculture pedologist, produced reconnaissance soil maps of all the western Minnesota counties adjacent to the Red River of the North. Nikiforoff's soil maps accurately detail shoreline complexes of Lake Agassiz. Nikiforoff's contribution to an understanding of regional Pleistocene history was his alternate interpretation of the history of Lake Agassiz (Nikiforoff, 1947).

The glacial history of Minnesota has been summarized by H. E. Wright, Jr., in The Quaternary of the United States (Wright and Ruhe, 1965). A more recent summary in The Geology of Minnesota (Wright, 1972) uses a detailed stratigraphic study in southwestern Minnesota by Charles L. Matsch (1971),

Several stratigraphic studies of Pleistocene sediments have been conducted in the region. Mark M. Fenton has developed a stratigraphic framework for Quaternary sediments **in** southeastern Manitoba (Fenton, 1974). Similar studies have been done in northeastern North Dakota (Salomon, in press), the Red Lake River area of Minnesota (Harris, 1973), and the Red River Valley of North Dakota and Minnesota (Harris and others, 1974).

Concurrent with this study similar studies are being conducted. by Donald K. Sackreiter (in preparation) in the south half of the subsurface study area, and Howard C. Hobbs (in preparation), in northeastern North Dakota. A detailed study of the stratigraphy of Lake Agassiz sediments is being conducted by B. Michael Arndt (in preparation).

Field Methods

Surface geology

Mapping of the surface geology began in the spring of 1973 and was completed by the fall of 1974. County highway maps, scale 1:126 720 (1 inch: 2 miles), from the Minnesota Department of Highways were used as base maps for the field study. Base maps for the preliminary and final versions of the surficials maps were United States Geological Survey 1° by 2° topographic maps, scale 1;250 000 (1 inch: 3.95 miles).

Stereoscopic aerial photographs were used *to* produce field maps of the study area. For the western half of the Grand Forks-Bemidji area Army Map Service aerial photographs, scale 1:58 940 (1 inch: 0.93 miles), flown in August 1952, were used. For the eastern half of the area aerial photographs from Mark Hurd Aerial Surveys, Inc., scale 1:86 400 (1 inch: 1.36 miles), flown in April 1969, were used. Soil maps were used where available to aid in lithologic interpretations of aerial photographs (Nikiforoff, 1933).

Field maps were field checked by driving most of the passable roads in the study area. Lithologic information was gathered by sampling surface exposures or by the use of a hand auger in areas of poor exposures.

A preliminary map was compiled at the final scale (1:250 000), and problem areas were again field checked.

Subsurface geology

Subsurface data used in this study were in part gathered in connection with a detailed study of the Pleistocene stratigraphy

exposed along the Red Lake River, Minnesota (Harris, 1973). In addition to the measured sections in river cutbanks, five shallow power-auger test holes, as much as 24 metres (80 feet) deep, were drilled in 1971. These data were supplemented by six deep power-auger test holes, as much as 46 metres (150 feet) deep, and six shallow power-auger test holes in 1973 and 1974.

The field procedure used in all test drilling included the onsite description and collection of samples from power-auger cuttings. In general either al metre (3 foot) or. a 1.5 metre (5 foot) sample interval was used.

Laboratory Methods

Textural data

All surface and subsurface samples were analyzed for content of sand (2 mm to $1/16$ mm), silt (1/16 mm to $1/256$ mm), and clay (less than 1/256 mm). Textural analysis was conducted in the North Dakota Geological Survey sediment laboratory. Results were compiled, and means and standard deviations were calculated for each lithostratigraphic unit present in the study area. Tabulation of the results of the textural analysis for surface data is in Appendix I. Tabulation of the results of the textural analysis for subsurface data is in Appendix II.

Coarse-sand lithology

All surface and subsurface samples were analyzed for content of crystalline- and metamorphic-rock types, carbonate-rock types, and shale rock fragments in the coarse-sand (1 mm to 2 mm) fraction. The coarsesand fraction was separated from the samples during textural analysis.

The proportion of each lithologic group was determined by studying individual sample fractions under a binocular microscope. Other material present in minor amounts in the coarse-sand fraction included lignite, wood, shell fragments, chert, sandstone, secondary carbonates, iron oxides, gypsum, and pyrite. The coarse-sand fractions counted ranged from 100 to 1500 grains but averaged about 400 grains. Tabulation of the results of coarse-sand lithology analysis for surface data is in Appendix I. Tabulation of the results of the coarse-sand lithology analysis for subsurface data is in Appendix II,

SURFACE GEOLOGY

Surface Map

The surface geology of the study area is presented on the map "Surficial sediments of the Grand Forks-Bemidji Area," (Plate 1). This map is a lithologic and genetic map of the surface sediments in the study area *at* a reconnaissance scale (1:250 000). The map units are descriptive lithologic units that are genetically interpreted in the legend.

Lithologic map units

Eight lithologic map units were used on the map. These include silt and clay; sand, silt, and clay; sand; sand and gravel; pebbleloam; sand overlying pebble-loam; peat overlying pebble-loam; and peat. These map units were chosen because they best describe the range of sediments encountered in the area mapped. Lithologic units less than 0.6 square kilometres (0.25 square miles) in area were not mapped. In general a minimum thickness of 1 metre (3 feet) was used as a mapping cutoff for these map units.

Silt and clay

This map unit indicates the location of material ranging from silty clay to clayey silt. The silt and clay contains little or no organic debris and is weakly to strongly laminated. This map unit is the dominant unit in the western part of the map. The silt and clay

is interpreted to be offshore lake sediment deposited in the Lake Agassiz basin.

Sand, silt, and clay

The map unit indicates the location of variable amounts of interbedded sand, silt, and clay, and organic debris. It is generally poorly to moderately laminated. The sand, silt, and clay was deposited as fluvial overbank sediment in modern floodplains.

Sand

This map unit indicates the location of material that is mainly medium- to coarse-grained sand. The sand contains little or no organic debris and is variably bedded. In the western and northern parts of the study area, the sand is mostly of lacustrine origin and is associated with the Lake Agassiz shoreline complex. In the southern and eastern parts of the study area, the sand is mostly fluvial meltwater deposits. Small patches of eolian sand are present in the Fertile, Minnesota, area.

Sand and gravel

This map unit indicates the location of material ranging from sandy gravel to gravelly sand. It is variably bedded and sorted, and contains little or no organic debris. In the western and northern parts of the study area the sand and gravel is moderately to well sorted and is found in curved ridges associated with the Lake Agassiz shoreline complex. In the central and eastern parts of the study area the sand and gravel is generally poorly sorted and interpreted to be esker complexes and other outwash deposits.

Pebble-loam

This map unit indicates the location of a mixture of sand, $silt$, clay, pebbles, cobbles, and boulders in varying proportions. Pebbleloam is unbedded and unsorted. Pebble-loam is present in the southcentral and northeastern parts of the study area. Here the topography ranges from undulating to hummocky due to the collapse of the glacial sediment. Pebble-loam is also present in the north-central part of the area where the topography is flat to undulating. Here the topography has been modified by the action of the waters of Lake Agassiz.

Sand over pebble-loam

This map unit indicates the location of variable amounts of sand overlying pebble-loam. The overlying sand is as much as a metre $(3$ feet) thick and covers more than one half of the surface area of the map unit. Sand overlying pebble-loam is present mainly in the western and north-central parts of the study area. Its occurrence is generally restricted to the shoreline complex of Lake Agassiz.

Peat over pebble-loam

This map unit indicates the location of variable amounts of peat overlying pebble-loam. The overlying peat is as much as a metre (3 feet) thick and covers more than one half of the surface area of the map unit. Peat overlying pebble-loam is presently mainly in the poorly drained lowlands in the north-central part of the study area.

Peat

This map unit indicates the location of dark brown or black partially decomposed and disintegrated organic material. Peat is

deposited in shallow stagnant bogs and marshes. Accumulations of peat are found in poorly drained areas throughout the study area. They are most commonly found in the northern one-third and eastern one-half of the study area.

Genetic map units

Seven genetic map units are used on the surface map to indicate the interpreted origin of the descriptive map units. These units represent the results of the action of four geologic processes: glacial, fluvial, lacustrine, and eolian.

Alluvial sediment

This map unit indicates the location of interbedded sand, silt, clay and organic debris deposited as fluvial overbank sediment in modern floodplains. Alluvial sediment is poorly to moderately laminated.

Eolian sediment

This map unit indicates the location of wind-blown deposits of fine-grained sand. Areas of eolian deposits commonly contain stabilized dunes with some active blowouts.

Bog sediment

This map unit indicates the location of highly organic deposits accumulated in the shallow bogs and marshes common in poorly drained areas.

Lacustrine shoreline sediment

This map unit indicates the location of nearshore sediments and shoreline sediments. The topography of the lacustrine shoreline

sediment is generally flat but includes curved ridges. The generally flat areas contain sand and gravel or sand overlying pebble-loam. The sediments are variably bedded and are moderately to well sorted. The curved ridges contain sand and gravel. They are variably bedded and are moderately to well sorted.

Lacustrine offshore sediment

This map unit indicates the location of lacustrine sediment deposited in an offshore environment. The resulting topography has very low relief. The sediment deposited in this environment is moderately to strongly laminated silty clay and clayey silt.

Glaciofluvial sediment

This map unit indicates the location of fluvial bedload sediment deposited by glacial meltwater in or near glacial ice. The glaciofluvial sediment includes eskers, collapsed esker complexes, outwash plains, and other meltwater deposits. Glaciofluvial sediment is largely sandy gravel or gravelly sand. It is variably bedded and variably sorted.

Glacial sediment

This map unit indicates the location of sediment deposited by glacial ice. The sediment was deposited directly by active glacial ice or by the melting of stagnant ice. The resulting topography is flat, undulating, rolling, or hummocky. The glacial sediment consists of pebble-loam with small amounts of sand and gravel.

Geomorphology

A variety of geomorphic features is present in the Grand Forks-Bemidji area. The study area is divided into four main provinces based on the occurrence of these features (Figure 2). The Lake-Plain Province dominates the western edge of the study area. The Shoreline-Complex Province is present in the western area and dominates the northern part of the study area. The Glacial-Upland Province dominates the southcentral and southeastern parts of the study area. The Outwash Province is present in the southeastern part of the study area.

The Lake-Plain. Province

The Lake-Plain Province is a low-relief plain. It is largely underlain by laminated silt and clay deposits in the Lake Agassiz basin (Figure 2). Geomorphological features in the province are subtle and best seen on aerial photographs, The most striking feature in the Lake-Plain Province is the flat lake plain itself. Other geomorphological features present have an average relief of about a metre. Some of the other features present include compaction ridges, intersecting lineations, and stream channels.

A compaction ridge may be seen in the Beltrami area of Polk County (Figure 3). The discontinuous ridge follows a sinuous course from southeast to northwest and is as much as 2 metres (7 feet) higher than the surrounding lake plain. The ridge is thought to be cored by bedload sediments of an ancestral Sand Hill River. These sediments were deposited before Lake Agassiz flooded for the last time. A veneer of lake sediment covers most of the ridge.

Fig. 2. Geomorphologic provinces of the Grand Forks-Bemidji

Area.

- 1. Lake-Plain Province
2. Shoreline-Complex Pr
- 2. Shoreline-Complex Province
3. Glacial-Upland Province
- 3. Glacial-Upland Province
- 4. Outwash Province

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Fig. 3. Beltrami compaction ridge, Lake-Plain Province.

T.148N. -T.147 N.

 $18\,$

Intersecting lineations are commonly seen on the flat surface of the Lake-Plain Province, These lineations are low-relief ridges and grooves. Usually the lineations cannot be seen on the ground but are easily seen on aerial photographs (Figure 4). The lineations show a preferred northwest-southeast orientation. The intersecting lineations have been discussed by several authors since they were first observed by Horberg (1951). The most probable hypothesis is that the lineations are drag marks formed by floating lake ice during a shallow stage of Lake Agassiz (Clayton and others, 1965).

The Shoreline-Complex Province

The Shoreline-Complex Province (Figure 2) is a gently rolling to flat surface that has been shaped by lacustrine beach and nearbeach processes. The main geomorphologic features present in this province are beach ridges and wave-cut scarps. Other features include sand spits, peat bogs, lake basins, and river trenches.

The dominant features in the province are beach ridges (Figure 5). These features are curved ridges that rise as much as 5 metres (16 feet) above the surrounding land surface. The sediment in the ridges is moderately *to* well sorted sand and gravel. It is ripple cross-bedded to flat bedded. The beach ridges were formed at the margin of the glacial Lake Agassiz. The generally flat areas surrounding the beach ridges are composed of sand, sand overlying glacial sediment, and glacial sediment. The action of waves has reduced the relief on the. glacial sediment exposed in the area, and a residual layer of stones on the surface is common.

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Fig. 4. Intersecting lineations, Lake-Plain Province.

T.152 N. ÷. $T.151 N.$

Fig. 5. Beach ridges, Shoreline-Complex Province.
R.44W.

T.149N.
T140N. T.148N.

Wave-cut scarps are present in the Shoreline-Complex Province. The most prominent and persistent wave-cut scarp is the Campbell scarp (figure 6). This scarp trends from north to south through the province and for most of its length is cut into glacial sediment. The Campbell scarp marked the shoreline of Lake Agassiz at the Campbell level. The Herman scarp, another prominent scarp, marks the boundary between the Shoreline-Complex Province and the Glacial-Upland Province. It is best developed in the north-central and northeastern parts of the study area, where it is as much as *30* metres (100 feet) high. The Herman scarp marked the shoreline of Lake Agassiz at the Herman level.

The Glacial-Upland Province

The Glacial-Upland Province is the largest province in the study area (Figure 2). It is an undulating to hummocky surface of collapsed glacial sediment. A variety of geomorphic features are present in the province. They include eskers, esker complex, meltwater channels, hummocky glacial sediment, lake basins, and scarps marking glacial-sediment margins,

Eskers and esker complexes are the most striking features in the Glacial-Upland Province (Figure 7). They are sinuous, discontinuous ridges and complexes of ridges. The sediment in the ridges is generally poorly sorted and variably bedded sand and gravel. The sediment contains flat-bedded units and both large- and small-scale cross-bedded units. These ridges consist of fluvial bedload sediment deposited by meltwater streams that flowed in or on the glacial ice. Esker complexes are scroll-like aggregates of eskers as much as 8 kilometres (5 miles) wide. Eskers and esker complexes are variably overlain by collapsed glacial sediment.

Fig. 6. Campbell scarp, Shoreline-Complex Province.

T.151 N.
T.150 N.

Fig. 7. Eskers and esker complex, Glacial-Upland Province. Photography by **Mark** Hurd Aerial Surveys.

Meltwater channels are common throughout the province. They are usually associated with the discontinuous course of eskers and esker complexes. Most meltwater channels have been overridden and. thus are poorly defined.

The most obvious meltwater channel in the province is the McIntosh Channel. It trends south and west from Trail to Fertile in Polk County (Figure 8). This channel averages about 1.6 kilometres (1 mile) wide over its 60 kilometre (36 mile) length. The average depth of the channel is about 10 metres $(33$ feet). The history and origin of the McIntosh Channel are discussed in detail by Moran and Clayton (in preparation). Moran and Clayton suggest that the channel was formed by glacial meltwater flowing north into Lake Koochiching, an early lake in the Agassiz basin. Later the channel drained meltwater south from Lake Koochiching into Lake Climax, another early lake in the Agassiz basin.

A discontinuous, generally obscure scarp marks the glacialsediment margin at the boundary between the Glacial-Upland Province and the adjacent Outwash Province. A well developed section of this scarp is located about 8 kilometres (5 miles) northwest of Bagley in Clearwater County (Figure 9). The glacial sediment to the north of the scarp stands 15 metres (50 feet) above the adjacent outwash *to* the south. An esker complex trends southward toward the margin of the glacial sediment. At the glacial-sediment margin the esker complex was drained by a meltwater channel in the outwash sediment.

The relationship of the glacial-sediment margin to the areal distribution of lithostratigraphic units will be discussed in the stratigraphic interpretations section.

Fig. 8. McIntosh Channel, Glacial-Upland Province. Photography by Mark Hurd Aerial Surveys.

T.148 N.
T.147 N.

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Fig. 9. Well developed scarp along the glacial-sediment margin near Bagley, Glacial-Upland Province. Photograph by Mark Hurd Aerial Surveys.

The Outwash Province

The Outwash Province is a flat to undulating surface of sand and gravel in the south-central and southeastern parts of the Grand Forks-Bemidji area (Figure 2). The major geomorphologic features present in this province are outwash plains, meltwater channels, eskers, and esker complexes.

The outwash plains are generally flat areas of sand and gravel (Figure 10). They were deposited by glacial meltwater streams draining ice-marginal positions. The sediments in the outwash plains are variably bedded and moderately to well sorted. Bedding present includes mainly large and small scale cross-bedded units.

The meltwater channels, eskers, and esker complexes present in this province (Figure 10) are similar to those previously discussed in the Glacial-Upland Province. They consist of fluvial bedload sedimerrt deposited by meltwater streams draining ice-marginal positions. The sediment is generally poorly to moderately sorted and variably bedded. The bedding ranges from flat-bedded units to large- and small-scale cross-bedded units.

Fig. 10. Outwash plain, Outwash Province. Photography by Mark Hurd Aerial Surveys.

SUBSURFACE GEOLOGY

The subsurface geology of the Grand Forks-Bemidji area and an area of equal size immediately to the south was studied. The subsurface study area and the location of the data points are shown in Figure 11.

Lithostratigraphic Units

Seven lithostratigraphic units are recognized in test holes and surface exposures in the subsurface study area. There are six units of glacial origin and one unit of lacustrine origin. These lithostratigraphic units were characterized by their texture, coarsesand petrology, and stratigraphic position (Table 1). Samples from 16 power-auger test holes, two composite sections compiled from previous drilling programs, and two representative measured sections in the Red Lake River trench were used to characterize the lithostratigraphic units. Descriptive logs of all test holes, composite sections, and Red Lake River sections are contained in Appendix III. Tabulated textural and coarse-sand petrology data are presented in Appendix II. In addition to the seven lithostratigraphic units recognized in the test holes, five additional lithostratigraphic formations, recognized in earlier studies, are present in the subsurface in the western part of the study area. These lithostratigraphic units will be discussed from the oldest to youngest with respect to their distinguishing characteristics, source area, age, and correlation with the stratigraphic units of other workers in adjacent areas.

Fig. 11. Location of data points in area of subsurface study.

Table.1. Summary of textural and coarse-sand petrology data for lithostratigraphic units.

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Unnamed unit 1

This unit consists of unbedded glacial sediment. It is pebbly, . silty clay-loam.

Two metres (7 feet) of this unit is present in test hole N-1916 (Appendix III). Here the unit is composed of an average of 20% sand, 46% silt, and 34% clay. The coarse-sand fraction contains an average of 61% crystalline- and metamorphic-rock fragments, 38% carbonate-rock fragments, and no shale fragments (Table l).

This unit is one of two unnamed units which are present beneath the Marcoux Formation. Unnamed unit 1 underlies unnamed unit 2 and is the oldest lithostratigraphic unit recognized in test holes made during this study. Unnamed unit l contains an average of 20% sand, less than any other units present except the Huot Formation. The amount of carbonate-rock fragments present in the coarse-sand fraction is greater than any other lithostratigraphic unit encountered in the study area. It contains an average of less than one percent shale fragments (Table 1).

The source area of this unit is unknown. The relatively high carbonate-rock fragment content suggests a north-northwesterly source area.

The age of unnamed unit 1 is unknown, but stratigraphic position suggests it is pre-Wisconsinan or Early Wisconsinan in age.

Unnamed unit 2

This unit consists of unbedded glacial sediment. It is pebbly clay-loam.

Unnamed unit 2 is found in test holes N-1915 and N-1916 (Appendix III). The unit averages about 4.5 metres (15 feet) thick. This unit is composed of an average of 39% sand, 32% silt, and 28% clay. The coarse-sand fraction contains an average of 70% crystallineand metamorphic-rock fragments, 29% carbonate-rock fragments, and 1% shale fragments (Table 1).

Unnamed unit 2 is significantly sandier than the underlying unnamed unit land significantly less sandy than the overlying Marcoux Formation. The coarse-sand fraction contains fewer carbonaterock fragments than unnamed unit land more than the Marcoux Formation. It contains an average of 1% shale (Table 1).

The source area of unnamed unit 2 is unknown. The ratio of crystalline- and metamorphic-rock fragments to carbonate-rock fragments suggests a north-northwesterly source area.

The age of unnamed unit 2 is unknown, but stratigraphic position suggests it is pre-Wisconsinan or Early Wisconsinan in age.

Marcoux Formation

The Marcoux Formation consists largely of unbedded glacial sediment. It is pebble-loam. A formal description of the unit is contained in Late Quaternary stratigraphic nomenclature, Red River Valley, North Dakota and Minnesota (Harris and others, 1974).

This unit is present in 13 test holes located throughout the area studied (Appendix III). Test holes rarely reached the base of the Marcoux Formation. An average thickness of 15 metres (50 feet) and a maximum thickness of 44 metres (146 feet) was encountered.

The Marcoux Formation is composed of an average of $50%$ sand, 35% silt, and 16% clay. The coarse-sand fraction contains an average

of 82% crystalline- and metamorphic-rock fragments, 16% carbonate-rock fragments, and 2% shale fragments (Table 1).

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The Marcoux Formation includes the surface glacial sediment throughout the southeastern part of the study area (Figure 12). In most of the subsurface the Marcoux Formation is overlain by the St. Hilaire Formation. In the Red Lake Falls area the St. Hilaire Formation is absent and the Marcoux is overlain by the Red Lake Falls Formation (Plate 2, Panel 1). The Marcoux Formation overlies unnamed unit 2.

The characteristics that are most useful in identifying the Marcoux Formation are its relatively high sand content, abundant crystalline- and metamorphic-rock fragments in the coarse-sand fraction, and low shale content. The average sand content is 50%, higher than any other unit present, The coarse-sand fraction contains 82% crystalline- and metamorphic-rock fragments, more than any other unit present (Table 1).

The Marcoux Formation is thought to have a north-northeastern source area because the coarse-sand fraction contains a large portion of Canadian Shield type rocks.

The age of the Marcoux Formation is unknown, but stratigraphic position suggests that it is pre-Wisconsinan or Early Wisconsinan in age.

The Marcoux Formation is correlatiod with the Vang Formation in northeastern North Dakota (Hobbs, in preparation). It is also correlated with the Hawk Creek Till (Matsch, 1971) in the Minnesota River Valley of southwestern Minnesota (Harris and others, 1974) (Table 2).

Fig. 12. Surface exposure of the lithostratigraphic units in the area of subsurface study.

 $94'30'W.$

Table 2. Correlation of lithostratigraphic units present in the subsurface study area with stratigraphic units of workers in some other areas.

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St. Hilaire Formation

The St. Hilaire Formation consists largely of unbedded glacial sediment. It is a pebble-loam. A formal description of the unit is contained in Late Quaternary stratigraphic nomenclature, Red River Valley, North Dakota and Minnesota (Harris and others, 1974).

This unit is present in nine test holes located in all but the southeastern part of the area studied (Appendix III). An average thickness of 11 metres (36 feet) was encountered.

The St. Hilaire Formation is composed of an average of 34% sand, 40% silt, and 26% clay. The coarse-sand fraction contains an average or 53% crystalline- and metamorphic-rock fragments, 24% carbonate-rock fragments, and 23% shale fragments (Table 1).

The St. Hilaire Formation contains the surface glacial sediment in a narrow band from the south-central part of the area studies, toward Bagley and southeastward toward Leech Lake (Figure 12). The St. Hi.laire Formation is overlain by the Red Lake Falls Fonnation throughout the northern and western parts of the area studied.

The characteristic most useful in distinguishing the St. Hilaire Formation from other units present is its high shale content. It contains an average of 23% shale fragments in the coarsesand fraction, more than any other unit (Table 1).

The St. Hilaire Formation is thought to have a westnorthwesterly source area. The shale fragments are thought to have come from the Pierre and Riding Mountain Formations in eastern North Dakota and southern Manitoba.

Based on stratigraphic position the St. Hilaire Formation is Wisconsinan or pre-Wisconsinan in age.

Fenton (1974) correlates his Stuartburn Till of southeastern Manitoba with the St. Hilaire. In northeastern North Dakota the St. Hilaire is correlated with the Gardar Formation (Hobbs, in preparation). The Granite Falls Till of southwestern Minnesota (Matsch, 1971) is thought to correlate with the St. Hilaire Formation (Table 2).

Red Lake Falls Formation

The Red Lake Falls Formation consists largely of unbedded glacial sediment. It is a pebble~loam. A formal description of the unit is contained in Late Quaternary stratigraphic nomenclature, Red River Valley, North Dakota and Minnesota (Harris and others, 1974).

This unit is present in 15 test holes located in all but the southeastern part of the area studied (Appendix III). An average thickness of 13 metres (42 feet) was encountered.

The Red Lake Falls Formation is composed of an average of 37% sand, 42% silt, and 21% clay. The coarse-sand fraction contains an average of 58% crystalline- and metamorphic-rock fragments (Table 1).

The Red Lake Falls Formation includes the surface glacial sediment throughout the central and northeastern parts of the area studied (Figure 12). In the northwestern part of the area it is overlain by the Huot Formation.

The characteristics of the Red Lake Falls Formation that distinguish it from other units are its stratigraphic position above the St. Hilaire Formation, its intermediate shale content (5%), and its loamy texture. Texturally the Red Lake Falls and St. Hilaire Formations are quite similar. The lower shale content of the Red Lake Falls easily distinguishes it from the St. Hilaire Formation. The

Red Lake Falls Formation is considerably sandier than the overlying Huot Formation (Table 1).

On the basis of coarse-sand lithology it.is thought that the Red Lake Falls Formation had a source area *to* the north-northwest.

The Red Lake Falls Formation is thought to be Wisconsinan in age because of its stratigraphic position beneath the Huot Formation and its extensive surface exposure (Figure 12).

In some outcrops along the Red Lake River and some test holes, two subdivisions can be distinguished within the Red Lake Falls Formation. This distinction cannot be made consistently. No subdivision of the Red Lake Falls Formation is made in this report. Where the distinction is made, in the Red Lake Falls area, the lower unit is somewhat sandier and the upper unit is considerably more shaley (Harris and others, 1974).

In northeastern North Dakota the upper part of the Red Lake Falls Formation is correlated with the Dahlen Formation (Hobbs, in preparation). In southeastern Manitoba the lower part of the Red Lake Falls Formation is correlated with the Senkiw Till and the upper part of the Red take Falls Formation with the Roseau Till by Fenton (1974). In southwestern Minnesota the New Ulm Till of Matsch (1971) is correlated with the Red Lake Falls Formation (Table 2).

Huot Formation

The Huot Formation consists largely of glacial sediment. It is slightly pebbly, silty clay-loam. A formal description of the unit is contained in Late Quaternary stratigraphic nomenclature, Red River Valley, North Dakota and Minnesota (Harris and others, 1974).

This unit is present in three test holes and one of the measured sections along the Red Lake River (Appendix III). An average thickness of 4 metres (14 feet) is present.

The Huot Formation is composed of an average of 8% sand, 34% silt, and 57% clay. The coarse~sand fraction contains an average of 59% crystalline- and metamorphic-rock fragments, 35% carbonate-rock fragments, and 6% shale fragments (Table 1).

In exposure along the Red Lake River the Huot Formation is underlain by the Wylie Formation, a laminated silt and clay of lacustrine origin (Harris and others, 1974). The Ruot Formation is the surface glacial sediment in a narrow band extending from the westcentral part of the study area northeast through the Red Lake Falls area (Figure 12).

The Huot Formation is easily identified by its stratigraphic position, texture, and coarse-sand lithology. It overlies the Red Lake Falls Formation and contains less sand and more clay than any other unit present in the area of study. The Huot contains more carbonate-rock than any of the other units present, except the Red Lake Falls Formation and unnamed unit 1 (Table 1).

The Huot Formation was deposited by latest Wisconsinan glacial ice that flowed south in the Red River Valley. The Edinburg Moraine, a thickening of the Huot Formation, marks the southern limit of the Huot Formation (Arndt, in preparation).

The Huot Formation is the lateral equivalent of the Falconer Formation in the northern Red River Valley (Harris and others, 1974) and northeastern North Dakota (Hobbs, in preparation}. In

southeastern Manitoba the Huot and Falconer Formations are correlated with the Marchand Till (Fenton, 1974) (Table 2).

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Sherack Formation

The Sherack Formation consists largely of laminated offshore lake sediment. It is composed of interbedded silt and clay. A formal description of the Sherack Formation is contained in Late Quaternary stratigraphic nomenclature, Red River Valley, North Dakota and Minnesota (Harris and others, 1974).

The Sherack Formation averages between 4.5 to 9 metres (15 to 30 feet) thick (Harris and others, 1974).

The Sherack Formation is easily distinguished from other lithostratigraphic units in the study area by its stratigraphic position and texture. In the southwestern part of the study area the Sherack is underlain by the Argusville Formation (Arndt, in preparation). In the northwestern part of the study area the Sherack is underlain by the Brenna Formation (Harris and others, 1974). The Argusville and Brenna Formations are massive, obscurely laminated clays of lacustrine origin. The Poplar River Formation occurs as scattered deposits of sand and gravel beneath the Sherack (Harris and others, 1974). The Sherack Formation is exposed on the surface throughout most of the western part of the study area, except where the Huot Formation is at the surface (Figure 12). The Sherack Formation is only wide-spread, conspicuously laminated lake sediment present in the study area. Its interbedded silt and clay easily distinguish it from other lithostratigraphic units encountered.

The Sherack Formation is latest Wisconsinan and Holocene in age. It was deposited as offshore sediment in the Agassiz basin between 10 000 and 9 500 BP (Harris and others, 1974}.

No correlation of Lake Agassiz sediments is made in this paper. A detailed stratigraphic study of the Agassiz basin is being conducted by B. Michael Arndt (in preparation).

Stratigraphic Interpretations

The stratigraphy of the area is shown in a generalized crosssection that extends across the area from northwest to southeast (Figure 13). In general, progressively younger lithostratigraphic units lap up and around a local highland in the southeastern part of the area. This highland consists mainly of glacial sediment of the Marcoux Formation.

Seven detailed cross-sections (Figure 14} have been constructed from test holes, composite sections, and Red Lake River sections (Appendix III). These cross-sections substantiate and elaborate the generalized stratigraphic framework. Three east-west cross sections (Panels 1, 2, and 3) and four north-south cross-sections (Panels 4, 5, 6, and 7) are contained in Plates $2, 3, 4,$ and 5 . These cross-sections contain the basic hole-to-hole correlations. As a matter of convention, sand and gravel units separating units of pebble-loam were placed in the formation containing the overlying pebble-loam.

The stratigraphic relationships of the Pleistocene sediments beneath the oldest lithostratigraphic unit encountered, unnamed unit 1, are unknown. Figure 13 shows that the total thickness of these sediments probably ranges from about 12 metres (40 feet) to about 60 metres

Fig. 13. Generalized stratigraphy of the subsurface study area.

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Fig. 14. Location and panel numbers of detailed crosssections.

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(200 feet). The elevation of the undulating bedrock surface (Moran, 1975} rises from about 600 feet in the northwest to about 1300 feet in the southeast.

Little is known about unnamed units 1 and 2. The intermediate content of crystalline-rock fragments and·the relatively high content of carbonate-rock fragments in the coarse-sand fraction (Table 1) suggests a north-northwesterly source area. It is thought that unnamed units 1 and 2 may be correlative with the glacial sediment of the Wadena drumlin field to the south of the subsurface study area (Sackreiter, in preparation).

The surface glacial sediment in the Itasca Highland in the southeastern part of the subsurface study area is in the Marcoux Formation. Drill holes in the Itasca Highland often encountered repetitive interval of sand and gravel separated by intervals of glacial sediment of the Marcoux Formation. Buried oxidized glacial sediments were encountered in Test holes N-1915 and N-1916 (Plate 4, Panel 5}. The repetitive buried oxidation zones and the sand and gravel units are interpreted as evidence of thrusting of glacial sediments near a glacial margin. The Itasca Highland is thought to be a morainic complex consisting largely of the Marcoux Formation.

The St. Hilaire Formation is thought to have been deposited by a single advance of glacial ice from the northwest. The sediment of the St. Hilaire Formation most often encountered in test holes is a homogeneous, shale-rich pebble-loam. Sand and graval members of the formation are not uncommon. They are most common near the southeastern limit of the formation in the study area (Figure 13). These sand and gravel members are thought to be meltwater sediments

deposited during the advance of the glacial ice that deposited the St. Hilaire Formation.

The Red Lake Falls Formation contains the surface glacial sediment throughout the central part of the subsurface study area (Figure 13). The extent of the glacial sediment of this formation in the subsurface study area is shown in Figure 13. In the Grand Forks-Bemidji area, the area of surface study, the southern limit of the glacial sediment of the Red Lake Falls Formation is marked by a discontinuous, generally obscure scarp. This discontinuous scarp marks the boundary between the Glacial-Upland Province, to the north, and the Outwash Province, to the south (Figure 2). A well developed scarp along the southern limit of the Red Lake Falls· glacial sediment is located about 6 kilometres (4 miles) north of Bagley (Figure 9). Meltwater discharged along this glacial margin flowed into a major meltwater channel that followed a southwesterly course along the glacial margin. The present day Clearwater River flows in this channel in the Bagley area.

The surface sand and gravel in the Bemidji area (Plate 1) is thought to be Red Lake Falls outwash sediment. This sand and gravel extends north and west under the surface glacial sediment of the Red Lake Falls Formation (Plate 3, Panel 4; Plate 4, Panel 5). Figure 15 shows the approximate extent of the overridden sand and gravel. In this area the thickness of the surface glacial sediment is variable, but generally ranges from about 3 metres (10 feet) to about 12 metres (40 feet). Overridden meltwater channels and esker complexes are common in this area, and sand and gravel is commonly exposed at the surface. The sand and gravel underlying the surface glacial sediment is

Fig. 15. Surface exposure of lithostratigraphic units in the area of subsurface study. Cross-hatched area indicates the approximate extent of the overridden sand and gravel of the Red Lake Falls Formation.

 $94'30'W$

thought to be overridden outwash sediment deposited along the margin of the glacier that deposited the Red Lake Falls Formation.

North and west of the overridden outwash sediments, the glacial sediments of the Red Lake Falls Formation may be divided into·an upper shalier member and a lower sandier member in some test holes. This distinction is noted on the detailed cross-sections where it can be made (Plate 2, Panel 1; Plate 4, Panel 5, Panel 6; Plate 5, Panel 7). It is tempting to correlate the lower sandier member with the overridden outwash sediments, and the upper shalier member with the surface glacial sediment of the Red Lake Falls Formation. However, the distinction between the upper and lower members cannot be consistently made in the subsurface, and the surface data indicates the presence of a low-shale surface glacial sediment throughout the area of exposure of the Red Lake Falls Formation. For the purposes of this report the Red Lake Falls Formation remains undifferentiated. More detailed stratigraphy is needed *to* fully understand the relationship between the upper and lower members of the Red Lake Falls Formation.

A moderately to well sorted, fine- to medium-grained sand is found beneath the glacial sediment of the Red Lake Falls Formation in some places in the northern part of the study area. This sand is present in test holes N-1011 and N-1914 and in the McIntosh, North composite section (Plate 2, Panel 1). A similar sand is present beneath 1.5 metres (5 feet) to 3 metres (10 feet) of Red Lake Falls glacial sediment in a borrow pit (SW₄, SW₄, SE₄ section 17, T-150N, R-34W) and in numerous poorer exposures in several surrounding townships (T-150N, R-33W; T-150N, R-34W; T-150N, R-35W; T-150N, R-36W). A similar sand is present at the surface in sections 10, 11, 15, and

15 of T-150N, R-36W and along the southeastern and eastern margins of Lower Red Lake (Plate 1). Location of exposures and detailed distribution of the aand on tribal lands of the Red Lake Indian Reservation was not possible because access to the reservation to field check airphoto . maps could not be gained. This apparently widespread sand unit is thought to be lacustrine shoreline sand associated, at least in part, with ponding of meltwater in the Agassiz basin before the deposition of the Red Lake Falls Formation.

QUATERNARY HISTORY

The Quaternary history of the area of study will be discussed event by event from oldest to youngest.

Glaciers advanced and retreated over the study area an unknown number of times, depositing the glacial sediment that probably occurs beneath the oldest glacial sediment present in test borings, unnamed unit 1 (Figure 13).

Glacial ice advanced over the study area and retreated twice, depositing the glacial sediment of unnamed units 1 and 2 (Figure 16). This glacial ice is thought to have flowed out of the Winnipeg lowland (Figure 17). The coarse-sand petrology of these units suggests a· north-northwesterly source area (Table 1).

The glacial ice of the Rainy and Superior Lobes advanced over the study area from the northeast (Figure 17) and retreated, depositing the Marcoux Formation (Figure 16). The coarse-sand petrology of the Marcoux Formation (Table 1) suggests a northeasterly source area. How far south this ice lobe extended is not known, but it is thought to have extended at least to southwestern Minnesota. At its maximum extent the glacial ice of the Rainy and Superior Lobes is thought to have eroded the Wadena drumlin field in pre-existing glacial sediment. A recessional ice margin is thought to have formed the morainic complex in the Itasca Highland.

Glacial ice that deposited the St. Hilaire Formation advanced into the area of study during the Early Wisconsinan (Figure 16). The

Fig. 16. Schematic time-distance diagram showing periods of deposition of formations present in the study area.

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Fig. 17. Quaternary events in the study area and adjacent areas.

a. Pre-Wisconsinan advance of glacial ice from the Winnipeg lowland. Deposition of unnamed units 1 and 2.

b. Pre-Wisconsinan advance of Rainy and Superior lobes from the northeast. Deposition of the Marcoux Formation (study area), Vang Formation {northeastern North Dakota), and Hawk Creek Till (southwestern Minnesota). Lines show position and orientation of Wadena drumlin field.

c. Early Wisconsinan advance of the Napoleon Glaciation from the northwest. Deposition of the St. Hilaire Formation (study area), Gardar Formation (northeastern North Dakota), Stuartburn Till (southeastern Manitoba), and Granite Falls Till (southwestern Minnesota).

coarse-sand petrology of the St. Hilaire Formation (Table 1) suggests a northwesterly source area. This glacial event is thought to be the Napoleon Glaciation in North Dakota (Clayton, 1962). The glacial ice was divided into two lobes by the Itasca Highland (Figure 17), an area ranging in elevation from 1500 to 2000 feet in the southeastern part of the study area. One large lobe flowed through the western part of the study area and south an unknown distance, down the Red River Valley (Figure 17). The other lobe flowed to the southeast around the Itasca Highland. How far southeast this lobe extended is unknown. The glacial ice of the Napoleon advance retreated from the area.

Glacial ice that deposited the Red Lake Falls Formation advanced into the study area during the Wisconsinan (Figure 16). The coarse-sand petrology of the Red Lake Falls Formation (Table 1) suggests a northwesterly source area and this glacial event is thought to be the Lostwood Glaciation (Clayton, 1972) in North Dakota. This glacial ice was divided into two lobes by the Itasca Highland (Figure 18). One large lobe flowed south down the Red River Valley and is thought to have deposited the bulk of the surface glacial sediment in western Minnesota. The other lobe flowed southeast around the Itasca Highland (Figure 18). How far southeast this lobe extended is unknown. The southern edge of the surface exposure of the Red Lake Falls Formation (Figure 15) marks the approximate position of the glacial ice margin as it flowed around the Itasca Highland. The glacial ice of the Lostwood advance retreated from the study area at about 14 000 BP (Harris and others, 1974).

As the glacier margin retreated northward, a proglacial lake occupied the Agassiz basin (Figure 18). The lake sediment of the

Fig. 18. Quaternary events in the study area and adjacent areas.

a. Wisconsinan advance of the Lostwood Glaciation from the northwest. Deposition of the Red Lake Falls Formation (study area), Roseau and Senkiw Tills (southeastern Manitoba), Dahlen Formation (northeastern North Dakota) and New Ulm Till (southwestern Minnesota).

b. Early phase of Lake Agassiz (Late Wisconsinan), deposition of the Wylie Formation (northern study area), lower Argusville Formation (southern study area), and Grunthal Formation (southeastern Manitoba).

c. Late Wisconsinan Caledonia Advance from the north. Deposition of the Huot and Falconer Formations (study area and northeastern North Dakota) and Marchand Till (southeastern Manitoba).

Wylie Formation (Harris and others, 1974) and the lower part of the Argusville Formation (Arndt, in preparation) were deposited in the Agassiz basin (Figure 15).

The Late Wisconsinan glacial ice of the Caledonia Advance deposited the Huot and Falconer Formations in the northwestern part of the study area (Figure 16). The southern extent of the Caledonia advance (Figure 18) is marked by the Edinburg **Moraine.** Arndt (in preparation) has shown that the Huot Formation thickens at this position. Deposition of the Edinburg Moraine occurred at about 13 000 BP (Harris and others, 1974).

Proglacial ponding continued in the Agassiz basin during the Caledonia advance. This phase of Lake Agassiz has been called the Lockhart Phase (Moran and Clayton, in preparation). The lake sediment of the Brenna Formation (Harris and others, 1974) and the upper part of the Argusville Formation (Arndt, in preparation) was deposited in the Agassiz basin between about 13 000 and 11 000 BP (Figure 16; Figure 19).

With the retreat of the glacial ice that deposited the Huot and Falconer Formations, the proglacial lake occupying the Agassiz basin drained. This low-water phase of the Lake Agassiz has been called the Moorhead Phase of Lake Agassiz (Moran and Clayton, in preparation). During the Moorhead Phase, about 11 000 to 10 000 BP streams deposited the sediment of the Poplar River Formation (Harris and others, 1974) on the lake plain (Figure 16; Figure 19).

A readvance of glacial ice to the north blocked a northeastern outlet of Lake Agassiz at about 10 000 BP (Harris and others, 1974). The Agassiz basin was again flooded during the Emerson Phase

Fig. 19. Quaternary events in study area and adjacent areas.

a. Late Wisconsinan Lockhart Phase of Lake Agassiz. Deposition of the Brenna Formation (northern study area), upper Argusville Formation (southern study area), and lower Hazel Formation (southeastern Manitoba).

b, Latest Wisconsinan Moorhead Phase of Lake Agassiz. Deposition of the Poplar River Formation (study area) and upper Hazel Formation (southeastern Manitoba).

c. Early Holocene Emerson Phase of Lake Agassiz. Deposition of the Sherack Formation (study area) and upper Hazel Formation (southeastern Manitoba).

of Lake Agassiz (Moran and Clayton, in preparation). The sediment of the Sherack Formation was deposited in the Agassiz Basin (Figure 16; Figure 19). The glacial ice retreated and Lake Agassiz was drained at about 9500 BP (Harris and others, 1974).

SUMMARY

Unconsolidated sediments, ranging from pre-Wisconsinan or Wisconsinan to Holocene in age, completely cover the bedrock of the area studied. Included are glacial, glaciofluvial, lacustrine offshore, lacustrine shoreline, bog, eolian, and alluvial sediments. The distribution of surficial sediments in the Grand Forks-Bemidji area is shown on Plate 1.

The study area may be divided into four geomorphic provinces on the basis of the presence of different geomorphic features. The very flat Lake-Plain Province dominates the western part of the study area. The Shoreline Province extends from the southwest to the northeast across the study area. The Glacial-Upland Province dominates the south-central and eastern parts of the study area. The Outwash Province is present in the southeastern part of the study area. Figure 2 shows the location of the geomorphic provinces in the study area.

Seven lithostratigraphic units are recognized in test holes in the subsurface study area. The bulk of the material in six units is of glacial origin, and the material in one unit is largely of lacustrine origin. Unnamed unit 1, unnamed unit 2, the Marcoux Formation, and the St. Hilaire Formation are largely glacial sediment and pre-Wisconsinan or Early Wisconsinan in age. The Red Lake Falls Formation is largely glacial sediment and Wisconsinan in age. The Huot Formation is largely glacial sediment and is latest Wisconsinan

in age. The Sherack Formation is largely lacustrine sediment and is latest Wisconsinan and Holocene in age.

The lithostratigraphic units are recognized and distinguished by means of their texture, coarse-sand lithology, and stratigraphic position. The texture of the sediments is characterized by the portions of sand-, silt-, and clay-sized particles present. The coarsesand lithology is characterized by the proportions of crystallineand metamorphic-rock fragments, carbonate-rock fragments, and shale fragments present.

The younger units of glacial sediment lap upon and around a highland, cored by older glacial sediment, in the southeastern part of the subsurface study area. Figure 13 shows a generalized crosssection of the subsurface study area. Detatled cross-sections are included in Plates 2, 3, 4, and *5.* Figure 15 shows the surface exposure of the lithostratigraphic units. Lake sediment of Lake Agassiz is present in the western part of the subsurface study area. The lithostratigraphic units present are correlated with units of other workers in the regions (Table 2).

The Quaternary history of the study area consists of an unknown number of glacial advances and retreats depositing the glacial sediment that underlies the oldest lithostratigraphic unit observed, unnamed unit 1. Glaciers advanced across the area and retreated three times depositing the glacial sediment of unnamed unit 1, unnamed unit 2, and the Marcoux Formation. Glaciers advanced into the area, flowed around the Itasca Highland, and retreated twice, depositing the glacial sediment of the St. Hilaire and Red Lake Falls Formations. During the treat of the glacier that

deposited the Red Lake Falls Formation, a lake occupied the Agassiz basin. The sediments of the lower Argusville and the Wylie Formations were deposited in the basin. The glacial ice that deposited_ the Huot Formation advanced south down the Red River Valley. During its advance and retreat, water continued to be ponded in the Agassiz basin. The sediments of the upper Argusville and the Brenna Formations were deposited in the basin. Lake Agassiz drained and refilled. During the last phase of Lake Agassiz the Sherack Formation was deposited.

Suggestions for future workers include more detailed work in the area studied and similar studies in adjacent areas, particularly to the south and east, to obtain a more complete understanding of the Quaternary stratigraphy and history of the Upper Midwest.

In the area studied. deep test drilling to bedrock would provide information on the Quaternary sediments present below the oldest units described here.

Similar studies in adjacent areas would provide additional surface and subsurface information to increase our understanding of the areal extent and stratigraphic relationships of the lithostratigraphic units present. Correlations could be extended, and the Quaternary history of the region would be better understood.

APPENDIX I

TABULATED SURFACE DATA

TABULATED SURFACE DATA

This appendix contains a tabulated summary of laboratory analysis of surface samples (Table 3). Sample data are listed by townships. The townships are listed from west to east in each tier. The tiers are listed from north to south.

The table headings include N-number, location, surface elevation, sample depth, description, texture, and lithology. These headings are discussed below.

The N-number is a unique sample number in the North Dakota Geological Survey sample library.

The location of the sample site is given as the quarter of the quarter of the quarter (A is NE₄, B is NW₄, C is SW₄, and D is SE₄) of the section(S) in which it is located. The township in which the site is located is indicated by its township number (T) and range number (R).

The surface elevation of the sample site and sample depth are given in feet.

A description of the sample site is given in a nine character code of the form "ABBCCDEFF."

The "A" indicates that the sample is (1) or is not (0) from the upper unit present. The "B" indicates the thickness of the upper unit in feet. The "C" indicates the lithology of the upper unit (1 is gravel, 2 is sand, 3 is silt, 4 is clay, 5 is pebble-loam, 6 is peat). Silty sand would be listed as 32. The "D" is a slash that separates the upper unit from the lower unit. The "E" indicates that the sample is (1) or is not (0) from the lower unit. The " F'' indicates the lithology of the lower unit as in "C" above. A sample site with 3 feet

of sand overlying pebble-loam that was sampled would have a description of "O 3 2/1 5."

The texture of the sample is given as the percentage content of sand (SD) silt (SLT), and clay (CLY).

The lithology of the coarse-sand fraction of the sample is given as the percentage content of crystalline- and metamorphic-rock fragments (XTAL), carbonate-rock fragments (CO_3) and shale fragments (SH).

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APPENDIX II

TABULATED SUBSURFACE DATA

TABULATED SUBSURFACE DATA

This appendix contains a tabulated summary of laboratory analysis of subsurface samples (Table 4), a tabulated summary of occurrence and site averages of lithostratigraphic units (Table 5), and a tabulated summary of means, medians, modes, and standard deviations of site averages for lithostratigraphic units (Table 6).

The table headings for Table 4 include N-number, location, sample depth, texture, and lithology. These headings are discussed below.

The N-number is a unique sample number in the North Dakota Geological Survey sample library.

The location of the sample site is given as the quarter of the quarter of the quarter (A is NE₄, B is NW₄, C is SW₄, D is SE¹/₂) of the section(s) in which it is located. The township in which the site is located is indicated by its township number (T) and range number (R).

The elevation of the sample and the sample depth are given in feet.

The texture of the sample is **given** as the percentage content of sand (SD), silt (SLT), and clay (CLY).

The lithology of the coarse-sand fraction of the sample is given as the percentage content of crystalline- and metamorphic-rock fragments (XTAL), carbonate-rock fragments (CO_3) , and shale fragments **(SH).**

CCD 12 146 37

CCD 12 146 37

 $N =$

LOCATION $1/4$ S T R

 $\frac{1}{2}$

 $SAMP$. **DEPTH** TE X TURF

Table 5. Occurrence and site averages of lithostratigraphic $units.$

 $^\circ$ 117

 \bar{z}

 $\overline{}$

 \bar{z}

 \mathcal{A}

 $\Delta \phi$

 $\bar{\nu}$

 \sim \sim

 $\frac{1}{18}$

 $\hat{\mathcal{A}}$

 $\Delta \sim 10^7$

 $\hat{\boldsymbol{z}}$

 $\hat{\mathcal{A}}$

 $\sim 10^5$

l.

Table 6. Means, medians, modes, and standard deviations for site averages of lithostratigraphic units.

APPENDIX III

SECTION DESCRIPTIONS

This appendix includes section descriptions for power-auger test holes, Red Lake River sections, and composite sections used in the subsurface study area. A map indicating the location of these sections is included.

Fig. 20. Location of data points in area of subsurface study.

 $\sim 10^{-10}$

 $94'30'W.$

Power-auger Test Hole

NE4, NE4, NE4, sec. 30, T. 150 N., R. 44 W.

Elevation 1040 feet

Section description by K. L. Harris

line and carbonate pebbles; yellow brown; glacial sediment.

N-141

Power-auger Test Hole

NW4, NW4, N W4, sec. 25, T. 149 N., R. 44 W.

Elevation 1150 feet

Section description by K. L. Harris

Power-auger Test Hole

SE%, SE%, SE%, sec. 17, T. 147 N., R. 45 W.

Elevation 956 feet

Section description by K. L. Harris

N-146

Power-auger Test Hole

SE4, SE4, NE4, sec. 28, T. 145 N., R. 45 W.

Elevation 955 feet

Section description by K. L. Harris

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Twin Valley Dam Site

Composite Section

SW₄, sec. 26, T. 144 N., R. 44 W.

Elevation 1105 feet

Section description by K. L. Harris

Formation

 $0 - 24$

1300-1276

Pebble-loam; sandy, silty; crystalline and carbonate pebbles; minor

Red Lake River Section

SW4, NE4, SE4, sec. 18, T. 151 N., R. 43 W.

Elevation 1051 feet

Section description by K. L. Harris

cobbles, and boulders; lower contact not exposed; glacial sediment.

N-1011

Red Lake River Section

NW~, SW~, SE~, sec. 16, T. 151 N., R. 44 W.

Elevation 1018 feet

Section description by K. L. Harris

Power-auger Test Hole

s~, ~,~,sec. 18, T. 148 N., R. *³⁷*w.

Elevation 1465 feet

Section description by S. R. Moran, K. L. Harris, and D. K. Sackreiter

Power-auger Test Hole

SW₄, SW₄, SE₄, sec. 12, T. 146 N., R. 37 W.

Elevation 1570 feet

Section description by S. R. Moran, D. K. Sackreiter, and K. L. Harris

Pebble-loam; yellow brown shale pebbles present; glacial sediment.

Unnamed unit 2

140-150 1430-1420

Pebble-loam; reddish brown; crystalline, carbonate, and reddish volcanic pebbles present; glacial sediment.

N-1916

Power-auger Test Hole

NW₄, SW₄, SW₄, sec. 28, T. 144 N., R. 37 W.

Elevation 1720 feet

Section description by **S. R.** Moran, D. K. Sackreiter, and K. L. Harris

Unnamed unit 2

Pebble-loam; sandy; glacial sediment.

Unnamed unit 1

148-156 1572-1564

Pebble-loam; silty; very hard; stony; greenish gray; glacial sediment.

N-1917

Power-auger Test Hole

SW~, Svh, SEl,;, sec. 6, T. 144 N., R. 41 W.

Elevation 1225 feet

Section description by S. R. Moran, D. K. Sackreiter, and K. L. Harris

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Power-auger Test Hole

 \mathbf{r}

 \mathbb{R}^2

NW₄, SW₄, NW₄, sec. 36, T. 142 N., R. 44 W.

Elevation 1170 feet

Section description by S. R. Moran, D. K. Sackreiter, and K. L. Harris

 \bar{t}

 $\mathcal{L}^{\text{max}}_{\text{max}}$

Power-auger Test Hole

NW₄, SW₄, NW₄, sec. 29, T. 147 N., R. 42 W.

Elevation 1225 feet

Section description by S. R. Moran and K. L. Harris

N-1920A '

Power-auger Test Hole

NE%, NW4, NW4, sec. 11, T. 150 N., R. 42 W.

Elevation 1126 feet

Section description by **S.** R. Moran and K. L. Harris

N-2470

Power-auger Test Hole

 $NW^{1}x$, $NW^{2}x$, $NW^{2}x$, sec. 28, T. 148 N., R. 34 W.

Elevation 1490 feet

Section description by D. K. Sackreiter and B. M. Arndt

1440-1426 Pebble-loam; sandy; carbonate and crystalline pebbles dominant; shale pebbles present; gray; glacial sediment.

N-2471

Power-auger Test Hole

SW₄, SE₄, SE₄, sec. 34, T. 150 N., R. 31 W.

Elevation 1360 feet

Section description by D. K. Sackreiter and B. M. Arndt

Power-auger Test Role

NW₄, NE₄, NW₄, sec. 11, T. 146 N., R. 31 W.

Elevation 1350 feet

Section description by D. K. Sackreiter

137

St. Hilaire Formation

Marcoux Formation

64

1286

Pebble-loam; silty, sandy; glacial sediment.

N-2474

Power-auger Test Hole

NEZ, NEZ, SEZ, sec. 1, T. 141 N., R. 31 W.

Elevation 1450 feet

Section description by D. K. Sackreiter and K. L. Harris

Power-auger Test Hole

$SW4, SW4, NW4, sec. 15, T. 142 N., R. 34 W.$

Elevation 1580 feet

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Section description by D. K. Sackreiter and K. L. Harris

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Harris, K. PhD 1975

Plate 1

SURFICIAL SEDIMENTS OF THE GRAND FORKS - BEMIDJI AREA

MAP UNIT

DESCRIPTION

a metre of variable amounts of
lt, clay, and organic debris. a metre of well sorted, cross-bedded, a metre of peat overlying pebble-loam. a metre of peat. a metre of moderately to well bly bedded sand. n a metre of poorly sorted, generally and and gravel. n a metre of moderately to well sorted,
ed sand and overlying pebble-loam. n a metre of laminated silt and clay. n a metre of laminated sand, silt, n a metre of variably bedded, ed sand. n a metre of variably bedded, poorly d gravel. n a metre of variably bedded, variably
verlying pebble-loam. n a metre of unbedded, unsorted Pebble-loam contains variable amounts cobbles, pebbles, sand, silt, and clay.

MAP UNIT BOUNDARIES

Harrio K: Ph.D. 1975

SYMBOLS

HORIZONTAL SCALE

 $\frac{8 km}{5 mi}$