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STRATIGRAPHY OF THE INYAN KARA FORMATION (LOWER CRETACEOUS) IN THE VICINITY OF THE NESSON ANTICLINE,

NORTHWESTERN NORTH DAKOTA

by

Brad L. Wartman

Bachelor of Science, University of Kansas

A Thesis

Submitted to the Graduate Faculty

of the

University of North Dakota

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in partial fulfillment of the requirements

for the degree of

- Master of Science

Grand Forks, North Dakota

May 1983 GEOL. T1983 Wa66

This thesis submitted by Brad L. Wartman in partial fulfillment of he requirements for the degree of Master of Science from the University North Dakota is hereby approved by the Faculty Advisory Committee nder whom the work was done.

Alan M. brancara (chairperson)

Richard D. to From

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This thesis meets the standards for appearance and conforms to the yle and format requirements of the Graduate School of the University North Dakota, and is hereby approved.

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Dean of the Graquate School

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Signature Brack L. Wartman

Date March 3, 1983

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ABSTRACT

The Lower Cretaceous Inyan Kara Formation, at the base of the Dakota Group, consists of sandstone, siltstone, shale, and coal. The stratigraphy of the Inyan Kara and facies relationships within the formation and with the underlying Swift and overlying Skull Creek Formations in the vicinity of the Nesson Anticline, northwestern North Dakota, were evaluated using 11 cross-sections, 163 borehole geophysical logs, and 1 well core.

The Inyan Kara can be differentiated on well logs into three members. The basal member, "A", consists of sandstone, siltstone, shale, and coal. It is about 200 feet (60 m) to 400 feet (122 m) thick in the study area. The member is characterized by abrupt facies changes of sandstone units. Log patterns in "A" suggest that the lower portion of the member was deposited in a dominantly deltaic environment whereas the upper portion of "A" was deposited in a dominantly fluvial environment. The member is approximately equivalent to the Lakota Formation in the Black Hills of South Dakota.

The middle member, "B", consists of sandstone, siltstone, and shale. It is about 20 feet (6 m) to 150 feet (46 m) thick in the study area. The member is characterized by gradual facies changes. Log patterns ^{suggest} that member "B" was deposited in a marginal-marine environment. The member is approximately equivalent to the Fall River Formation in the Black Hills.

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The upper member, "C", consists of interbedded siltstone and shale. It is about 15 feet (4.5m) to 40 feet (12 m) thick in the study area. The member is characterized by a lack of facies changes. The extent, continuity, and lithologies in "C" suggest that the member was deposited in a shallow marine environment. The member is approximately equivalent to the lower portion of the Skull Creek Formation in the Black Hills.

Facies relationships suggest that the upper portion of the Swift Formation and the lower portion of member "A" represent a regressional sequence in the study area whereas the upper portion of member "A", members "B" and "C", and the Skull Creek Formation represent a transgressional sequence in the study area.

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Thickness trends of members "A", "B", and "C" suggest that the Williston Basin and the Little Knife Anticline were active structures during the deposition of the Inyan Kara, that the direction of regression was from east to west, and that the direction of transgression was from west to east in the study area.

Evidence for an unconformity between the Swift and Inyan Kara Formations, as proposed by previous workers, in the study area is possibly ambiguous and inconclusive. Additional evidence suggests that both conformable and unconformable contacts exist between the Swift and the Inyan Kara.

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INTRODUCTION

In 1979, the North Dakota Geological Survey, in a co-operative effort with the Department of Energy, Division of Geothermal Energy, began a multi-year study to assess the hydrothermal potential of regional aquifers in the state. Preliminary results of regional studies indicated that the Inyan Kara Formation was one of the most promising aquifers, with good potential for development as a low-temperature hydrothermal resource (Harris and others 1981). The present study was undertaken to provide a more detailed understanding of the formation in the northwestern part of the state.

Purpose of Study

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In this study, work on the Inyan Kara was directed towards five purposes:

- (1) To determine the lithology and thickness of the Inyan Kara Formation in northwestern North Dakota.
- (2) To evaluate the presence and extent of members within the Inyan Kara in northwestern North Dakota.
- (3) To interpret possible facies relationships of these members.
- (4) To evaluate the equivalence of members of the Inyan Kara in northwestern North Dakota to rock-stratigraphic units in the type area, the Black Hills of South Dakota.
- (5) To interpret possible depositional environments of members within the Inyan Kara in northwestern North Dakota.

Geologic Setting

Stratigraphy

In North Dakota, the Inyan Kara Formation is the basal unit in a series of clastics that form the Lower Cretaceous Dakota Group (Fig. 1). The formation is generally regarded to be unconformable with the underlying Swift Formation (Upper Jurassic) and in transitional contact with the overlying Skull Creek Formation (Hansen 1955, Anderson 1982).

The Inyan Kara consists of sandstone, shale, siltstone, and coal (Hansen 1955). Deposition of the formation is generally regarded to have been in environments varying from non-marine to marine (Butler 1981, Anderson 1982).

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Regional Structures

Structures in North Dakota are largely restricted to the western two-thirds of the state. The largest structure is the Williston Basin, which covers North Dakota and extends into South Dakota, Montana, Saskatchewan, and Manitoba (Gerhard and others, 1982). Regional dips in the Basin are towards the west-central portion of North Dakota.

Five major anticlines are present in the Williston Basin. The Billings Anticline, Little Knife Anticline, and Nesson Anticline trend north-south. The Cedar Creek Anticline and Antelope Anticline trend northwest-southeast (Fig. 2).

Figure 1. Stratigraphic column of the Dakota Group (modified from Bluemle and others 1981).

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Series	Group	Formation
U. Cretaceous	Colorado	Belle Fourche
		Mowry
		Newcastle
		Skull Creek
L. Cretaceous	Dakota	Inyan Kara
U. Jurassic		Swift

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Figure 2. Location of major anticlinal features in western North Dakota (modified from Gerhard and others 1982). Study area is shown by dashed line.



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Location of Study Area

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The area of study for this project covers approximately 7,524 square miles (19,487 square kilometers) in northwestern North Dakota (Figs. 2 and 3). The study area was chosen for two reasons:

- (1) The density of oil and gas wells drilled in the region is high, allowing for detailed formation correlation, by relatively close well spacing, on the order of a few miles.
- (2) The area contains several structures, including the Nesson, Little Knife, and Antelope Anticlines and the central portion of the Williston Basin, which may have had an influence on the deposition of the Inyan Kara.

Previous Work

The stratigraphic nomenclature of the Inyan Kara and its lateral equivalents has undergone considerable change since its inception in the mid-1800s. The review of nomenclature in this study will concentrate on the work done in South Dakota, primarily in the Black Hills, and in North Dakota. Figure 4 is a summary of the units currently accepted as equivalent to the Inyan Kara in South Dakota, Montana, Saskatchewan, and Manitoba. Figure 5 summarizes stratigraphic terminology for units equivalent to the Inyan Kara in South Dakota, and Figure 6 shows stratigraphic terminology for the Inyan Kara in North Dakota.

Figure 3. Location of study area in northwestern North Dakota.

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Figure 4.

Stratigraphic units equivalent to the Inyan Kara Formation in areas surrounding North Dakota (modified from Bluemle and others 1981).

	1				
Montana	South Dakota	North Dakota		Manitoba	Saskatchewan
Skull Creek	Skull Creek		Skull Creek	Ashville	Joli Fou
'Basal Colorado Silt'					
Dakota	Fall River	n part)			
Kootenai		l Group (ir	inyan Kara	Swan River	Cantuar
Lakota	Lakota	Dakota			· ·
Morrison	Morrison		Swift	Waskada	Vanguard

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Figure 5. Summary of nomenclatural history of units equivalent to the Inyan Kara Formation in South Dakota. :

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Meek and Hayden (1858)	Darton (1901,1909)	Russel (1928)	Rubey (1930)		RubeyGriesWaageRice(1930)(1954)(1959)(1976)		Waage (1959)		Rice (1976)	Bluemle and others (1981)
Unit no. 2 (Ft. Benton Gp.)	Ft. Benton Gp.	Graneros Shale	Skull Creek			Skuli Creek		Skull Creek	Skull <u>Creek</u> "Basal silt"	Skull Creek
	Dakota	Fall River		Fall River	part)	Fall River		Fail River	Fall River	Fall River
Basal no. 1 sand (Dakota)	Fuson	Fuson	ra Gp.	Fuson	Dakota Gp. (in	Fuson	ra Gp.		Fuson \	
	Lakota	Lakota	Inyan Kar	Lakota		Lakota	Inyan Ka	Lakota	Lakota	Lakota
Beulah Shale	Beulah Shale	Morrison	м	orrison	М	orrison	м	orrison	Morrison	Morrison

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Figure 6. Summary of nomenclatural history of the Inyan Kara Formation in North Dakota.

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Kline (1942) Hans		Hansen (1955)		Anderson nd Carlson (1966)	Rice (1976)	l E	Bluemle and others (1981)	This study										
Ft. Benton Gp.	Dakota Gp. (in part)	Skull Creek	Skull Ç Creek		Skull Creek	¢	Skull Creek	rt)	0	Skull Creek								
Dakota		Fall River G U) -dg Fuson B G C Fuson B G C Fuson B C B C C C C C C C C C C C C C C C C C	(in par	(in par	(in par	in par	in par	in par	(in par	Fall River	in par	Fall River	Fall River	in par		(in pa	на. Га	*B*
Fuson				Fuson	G Inyan G Kara		a Gp.	Kara	1									
Lakota		Lakota	Dakot	Lanoia	Lakota	Dakot		Dakot	Inyan	-A-								
Beulah Shale	Morrison		Swift		Swift		Swift		Swift									

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<u>South Dakota</u>.-- The first recorded reference to Cretaceous rocks in the Upper Midwest was made by Lewis and Clark in 1805 during their expedition to the Pacific Northwest (Schoon 1971). Near a campsite on the Missouri River, in what is now northeastern Nebraska, the expedition record described a cliff of buff-colored, massively-bedded sandstone that is now recognized as a part of the Dakota Formation.

In South Dakota, the first description of Cretaceous rocks was made by Meek and Hayden (1858), who were naturalists accompanying an Army expedition to the Black Hills, then a part of the Nebraska Territory. They described a sequence of light brown sandstones, variously colored shales, and beds of coal that they recognized as the basal Cretaceous strata in the Black Hills. Meek and Hayden correlated these strata to the "Basal Number One Sand" (Dakota Formation) that, earlier in the expedition, they had described and named in what is now northeastern Nebraska, near the town of Dakota City.

Newton and Jenny (1880) studied and described cycadophytes from the lower part of the Dakota Formation (of Meek and Hayden) and ascribed an age of Early Cretaceous to the strata.

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Ward (1894, 1899) restudied the cycadophytes of the Dakota of the Black Hills. He concluded that the Dakota (of Meek and Hayden) could be divided into a lower member of Early Cretaceous age and an upper member--equivalent to the true "Dakota"--of Late Cretaceous age.

Darton (1901, 1909) correlated rock units from the Front Range of the Rocky Mountains to the Black Hills. He considered Meek and Hayden's ^{Cak}ota Formation as consisting of three formations. He named his basal ^{Cormation} the Lakota, a sequence of coarse-grained, massive, and cross-

bedded sandstones, with local conglomerates, 200-300 feet (60-91 m) thick. He noted the presence of fossils (cycadophytes and saurian bones), and stated that, on the basis of sedimentary structures, such as local channeling, the basal contact of the Lakota with the underlying Morrison Formation was unconformable. Darton named his middle formation, a sequence of nonfossiliferous clay and fine-grained sandstone averaging 100 feet (30 m) thick, the Fuson. He named the uppermost unit, a sequence of nonfossiliferous, fine-grained sandstones averaging 100 feet (30 m) thick, the Dakota. Darton concluded that the contact between the Dakota and Fuson represented a time-stratigraphic boundary separating the rocks of Early Cretaceous age (Lakota and Fuson) from rocks of Late Cretaceous age (Dakota) in the Black Hills.

Russel (1928) described fossils, mostly cycadophytes, that had been discovered recently in Darton's Dakota Formation, and concluded that Darton's Dakota was actually of Early Cretaceous age. To help alleviate the confusion that was arising in the use of the term "Dakota", Russel suggested that the name "Fall River Formation" be substituted for Darton's Dakota Formation.

Rubey (1930) measured and described several sections of the Lakota-Fall River interval in and around the Black Hills. He concluded that Parton's Fuson Formation did not extend outside of the area directly surrounding the Black Hills, and that the Lakota-Fall River interval bould be combined together into the Inyan Kara Group.

Gries (1954) published the first subsurface correlations of Lower ^{ret}aceous rocks in South Dakota using borehole geophysical logs. He ^{grelated} units from the Black Hills to the eastern part of the state

d southward into northeastern Nebraska. Gries said that the Dakota ormation in eastern Nebraska and South Dakota was represented in the stern part of the state by the Newcastle Formation. He said that the oks of the Lakota-Fall River interval were not an extension of the kota Formation, but instead represented a separate period of position, with the sediments of the Dakota directly overlying the Fall ver in the eastern part of the state. Gries proposed that the term akota Group" be used to include all the strata of the Lakota-Newcastle terval.

Waage (1959) measured and described several sections of the Inyan a Group, including a few of Darton's original sections, in the Black Is. Waage concurred with Rubey's conclusion that Darton's Fuson mation was not correlatable outside of the Black Hills area. He ther suggested that the term "Fuson" be dropped altogether, and that strata of Darton's Fuson be included in the Lakota Formation. Waage luded that the boundary between his Lakota and Fall River Formations esented a "transgressive disconformity" marked by a zone of abundant rite pellets.

Rice (1976, 1977) constructed cross-sections and correlation charts retaceous and Paleocene rocks in South Dakota. He correlated the 8 of the Lakota-Fall River interval, including the Fuson Formation, modified Russel's classification by dropping the term "Inyan Kara ". Rice recognized the Lakota Formation as the basal Cretaceous stone unconformably overlying rocks of Jurassic age, the Fall River e first sandstone directly underlying the Skull Creek Formation, he intervening strata as within the Fuson Formation.

Schoon (1971) and Bluemle and others (1981) have published stratigraphic charts that have modified Waage's classification by dropping the term "Inyan Kara Group" and retaining the Lakota and Fall River.

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North Dakota.-- The nomenclatural history of Lower Cretaceous rocks in the state has generally paralleled the work done in South Dakota. The earliest classification in the state is that of Kline (1942), who in a general stratigraphic column of rock units in the Williston Basin, correlated Darton's Lakota, Fuson, and Dakota Formations with the basal Cretaceous sandstones and shales in the Basin.

Hansen (1955) modified Gries' Dakota Group to include the Mowry Formation, and renamed Kline's Dakota Formation the Fall River Formation. He constructed subsurface cross-sections of the Lakota-Greenhorn interval in the state using borehole geophysical logs, and included descriptions of drilling samples for the Lakota, Fuson, and Fall River Formations. Hansen described the Lakota Formation as a white to light gray, coarse to medium-grained sandstone with "shale streaks", and reaching a maximum thickness of 110 feet (34 m). He concluded that the Lakota was deposited as a sheet sandstone in a littoral environment and formed the basal facies of a transgressive sequence. Hansen described the Fuson Formation as a medium-gray to gray-black shale with lenses of very fine-grained, quartzose sandstone. Thickness of the Fuson was described as highly variable, with a maximum thickness of 80feet (24 m) in the western part of the state. He concluded that the Fuson was deposited in an epineritic environment to the south and west of the transgressive Lakota Formation. The uppermost unit, the Fall

River Formation, was described by Hansen as an interbedded sequence of light gray, coarse to fine-grained, quartzose sandstone and gray, massively bedded, silty shale. Thickness of the Fall River was given as 30-210 feet (9-64 m), with the thickest part of the formation occurring in the west and thinning to the east. Hansen separated the depositional environments of the Fall River into two areas:

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- (1) In the southern part of the state, the Fall River was deposited under unstable shelf conditions varying from neritic to littoral and fluvial.
- (2) In the northern part of the state, the Fall River was deposited in either a platform or deltaic environment, with transgression of the epicontinental Cretaceous sea from southwest to northeast in the state.

Anderson and Carlson (1966) followed Waage's classification by including the Fuson as a part of the Lakota Formation.

Rice (1976, 1977) published cross-sections and correlation charts of Cretaceous and Paleocene rocks extending across North Dakota, central and western South Dakota, and eastern Montana. He correlated the Lakota-Fall River interval, including the Fuson Formation, and modified Russel's classification by dropping the term "Inyan Kara Group". Rice recognized the Lakota Formation as the basal sandstone unconformably Overlying rocks of Jurassic age, the Fall River Formation as the first Sandstone directly underlying the Skull Creek Formation, and the Intervening strata as within the Fuson Formation.

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The North Dakota Geological Survey, after examination of hundreds of well logs, concluded that the Lakota-Fuson-Fall River subdivision was

ot applicable, and grouped all three formations into the Inyan Kara ormation (Bluemle and others 1981).

MATERIALS AND METHODS

Introduction

In the study area, information on the Inyan Kara was obtained from borehole geophysical logs (also known as well logs, wireline logs, or simply logs) and well core.

Interpretations of the extent and depositional environments of the Inyan Kara and its members were made from cross-sections, contour maps, and lithologic descriptions.

Borehole Geophysical Logs

Borehole geophysical logs were the most abundant data source available in the study area. Interpretations made from well logs were used to identify and correlate rock-stratigraphic boundaries and make qualitative assessments of lithology and depositional environments. In this study three types of well logs, the Gamma-Ray, Spontaneous Potential, and Focusing-Electrode (Laterolog), were used.

For each well log type, a short description of the parameters measured and applications of the log are given below. A more detailed description of each log type can be found in training manuals published by various logging companies (Schlumberger 1972, Dresser Atlas 1974, Welex 1978).
Gamma-Ray Log

The Gamma-Ray log measures the natural gamma-ray emissions produced by the decay of unstable isotopes of uranium, potassium, and thorium, which occur naturally in sediments. In this study, the Gamma-Ray log was used for two purposes:

- (1) To correlate rock-stratigraphic units.
- (2) To qualitatively estimate lithology for lithologic units greater than five feet (1.5 m) in thickness.

Spontaneous Potential Log

The Spontaneous Potential log measures the response of a formation to cation-anion movement along formation boundaries from an induced electrical current. In this study, The Spontaneous Potential log was used to aid the Gamma-Ray log in identifying and correlating sandstone units.

Focusing-Electrode Logs

The Focusing-Electrode logs are a suite of log types that measure formation resistivity at various distances into the formation by the projection of a current "beam". One type, the Laterolog, was chosen for use in this study because of its ability to resolve shallow and deep formation resistivities in thin beds (less than five feet (1.5 m) thick). In this study, two types of Laterologs were used:

(1) The Shallow Laterolog, used on the Dual Laterolog, measures the formation resistivity less than 6 feet (1.8 m) into the formation (invaded and flushed zones). (2) The Deep Laterolog, used on the Dual Laterolog, Laterolog-3, and Laterolog-7, measures the resistivity at 9 to 15 feet (2.7 to 4.5 m) into the formation (true formation resistivity).

The Laterolog was used in this study for three purposes:

- (1) To correlate rock-stratigraphic units.
- (2) As a qualitative indicator of formation lithology.
- (3) To provide a qualitative indication of formation permeability(Dual Laterolog only).

Lithologic Identification

Identification of lithologies on borehole geophysical logs is a matter of interpretation based on the response of a given lithology to the logging devices. In the Inyan Kara, a combination of the Gamma-Ray log and Dual Laterolog provided the best recognition of these lithologic types (Fig. 7):

- (1) Sandstone
 - (a) Gamma-Ray log
 - (i) Low radioactivity level. Sandstones with a low shale content are called "clean"; those with a relatively high shale content are called "dirty".
 - (b) Spontaneous Potential log
 - (i) Positive deflection of curve.
 - (c) Dual Laterolog
 - (i) Deep Laterolog resistivity is greater than the Shallow Laterolog. Separation of the Laterolog curves is generally wide.
 - (ii) Deep Laterolog resistivity tends to be low.

Figure 7. Gamma-Ray log and Dual Laterolog of the Inyan Kara Formation showing interpreted lithologies (NDGS 7226, Sec. 21, T. 155 N., R. 90 W.).



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(2) Siltstone

- (a) Gamma-Ray log
 - (i) Low radioactivity (not diagnostic).
- (b) Dual Laterolog
 - (i) Resistivity level variable, depending upon permeability, with values generally between those of sandstone and shale.
 - (ii) Narrow to moderate separation of shallow and deep

resistivity values.

(3) Shale

- (a) Gamma-Ray log
 - (i) Relatively high radioactivity level.
- (b) Dual Laterolog
 - (i) High, generally uniform resistivity.
 - (ii) Little or no separation of shallow and deep resistivity

(4) Coal

(a) Gamma-Ray log

values.

(i) Radioactivity level variable, with levels ranging from intermediary between sandstone and shale to greater than shale. 111

(b) Dual Laterolog

- (i) Resistivity level is greater than or equal to shale.
- (ii) Separation of shallow and deep resistivity values highly variable, depending upon the amount of fracturing

(not diagnostic).

Well Core

Only one well in the study area, the Matthew Iverson 1 (Sec. 1, T. 155 N., R. 95 W.), NDGS 165, had recovered core from the Inyan Kara. Two intervals of the formation were cored (Fig. 8):

(1) Interval 1: 4980 to 4930 feet (1517 to 1502 m).

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(2) Interval 2: 4647 to 4590 feet (1416 to 1399 m).

Drilling Samples

Although drilling samples were available for the Inyan Kara in the study area, the quality of these samples is questionable. Conversations with the personnel of oil and gas drilling rigs and wellsite geologists indicate that depth control and sample quality for formations younger than Jurassic are generally poor. Because of these discussions, it was decided not to include a description of any drilling samples in this study.

Cross-Sections

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Correlation of the Inyan Kara Formation and its members in the study area was made from nine cross-sections, using logs from 94 wells (Plate I). Six cross-sections were oriented east-west with a spacing between sections of approximately four townships (Plates II-IV). The remaining three cross-sections were oriented north-south (Plates V and VI). Additional wells were used, where available, between logs in a given section to aid in correlation. The top of the Mowry Formation,

Figure 8. Well log of NDGS 165 (Sec. 1, T. 155 N., R. 95 W.) showing cored interval of the Inyan Kara Formation in relationship to members "A", "B", and "C".

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which has a characteristic radioactivity pattern in the study area, was used as a stratigraphic datum.

Logs for the cross-sections were chosen using the following criteria:

- (1) One well per township along the line of section, with a minimum well spacing of 2 miles (3.2 km).
- (2) Resolution of log character as compared to other logs in the township.

In addition, two cross-sections were constructed for the Lone Butte and Stanley oil fields for detailed correlation of members and sandstone units of the Inyan Kara (Plate VII). Well spacing for these sections was less than one mile (1.6 km).

For both the regional and detailed cross-sections, correlation of stratigraphic units was based upon two criteria:

- (1) Similar stratigraphic position between wells.
- (2) Continuity of log character between wells.

The legal description of wells used in the cross-sections is given in Appendix A.

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Contour Maps

Logs from 583 wells in the western one-third of the state were used to provide data for regional structure contour and isopach maps of the Inyan Kara (Plates VIII and IX) using a combination of equal spacing and interpretive contouring techniques (Bishop 1960). Wells (Appendix B) were chosen on the basis of one well per township. Isopach maps of each member (Plates X-XII) of the Inyan Kara in the study area were made with logs from 163 wells. Wells (Appendix C) were chosen on the basis of one well per township.

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Lithologic Descriptions

Lithologic descriptions of the Inyan Kara and its members were based on:

- (1) Interpretations from well logs that provided qualitative information on general lithology.
- (2) Hand sample descriptions of well core that provided information on lithology and sedimentary structures.

A description of well core is given in Appendix D.

RESULTS

Inyan Kara Formation

Identification of formation boundaries

Lower contact.-- In the study area, four types of contacts mark the base of the Inyan Kara:

- (1) Abrupt transition from shale (Swift) to "clean" sandstone (Inyan Kara) (Well 45, Section E-E', Plate IV).
- (2) Abrupt transition from a relatively low-resistivity shale
 (Swift) to a high-resistivity shale (Inyan Kara) (Well 83,
 Section H-H', Plate V).
- (3) A gradational contact of interbedded sandstone and shale grading upwards into sandstone (Well 15, Section B--B', Plate II).
- (4) A "clean" sandstone (Inyan Kara) overlying a "dirty" sandstone(Swift) (Well 23, Section H-H', Plate V).

In most parts of the study area, the contact is of type (1). Contacts of types (1) and (2) are relatively simple to identify and correlate. In the few areas where types (3) and (4) are present, the contact is recognized by correlating log markers in the Swift Formation from the log in question with log markers in logs from nearby wells. Correlation of log markers is continued upsection until units can no longer be correlated; the contact is then chosen at the boundary between the correlatable (Swift) and non-correlatable (Inyan Kara) units.

<u>Upper contact</u>.-- The upper contact of the Inyan Kara with the Skull Creek Formation is a gradational zone of interbedded siltstone and shale grading upwards into shale in all parts of the study area. In this zone the contact was chosen at the point where a distinctive "sawtooth" pattern of alternating high and low resistivity values appears on the Laterolog (Fig. 7) going down section.

Structure

The surface of the Inyan Kara varies from -2000 feet (-610 m) below sea level in the north-central portion of the study area to -3386 feet (-1032 m) in the west-central portion (Plate VIII).

The largest structure in the study area is the Williston Basin, which provides a regional dip towards the west-central portion. The Williston Basin extends outside of the study area to eastern North Dakota, southwestern Manitoba, southeastern Saskatchewan, eastern Montana, and northwestern South Dakota (Gerhard and others 1982).

The most prominent structure in the study area is the Nesson Anticline, which reaches a maximum elevation of -2190 feet (-667 m) in southeastern Williams County. The Nesson plunges southward from the northern border of the study area, covering an area of approximately 750 square miles (1942 square kilometers). The western flank of the Nesson is steep, with a dip of approximately 900 feet/mile (170 m/km).

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The Little Knife Anticline, in the southwestern portion of the study area, trends north-south but produces little structural relief.

The Antelope Anticline trends northwest-southeast and intersects the southeastern corner of the Nesson Anticline.

There are three unnamed structures in the study area that occur in Dunn County. In the northwestern portion of the county, two depressions trend northeast-southwest. In the southeastern portion of the county, a positive structure trends northwest-southeast, plunging to the northwest.

Thickness

The Inyan Kara reaches a maximum thickness in the study area of 495 feet (151 m) in south-central Williams County and a minimum of 226 feet (69 m) in south-central Dunn County (Plate VIII).

The central and western portions of the study area are part of a broad region of sediment greater than 400 feet (122 m) thick that extends westward out of the state. Areas of formation thickness greater than 450 feet (137 m), in the central portion of the study area, occur west of the approximate axis of the Nesson Anticline. Two long, narrow extensions of sediment greater than 400 feet (122 m) thick project out of this region. One extends to the north along the axis of the Nesson Anticline. The other extends into the south-central portion of the study area.

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The northeastern and southeastern portions of the study area are dominated by linear zones of thick sediment (greater than 350 feet (107 m) thick). These zones are roughly oriented towards the west-central portion of the study area.

In the southwestern portion of the study area are two anomalous regions of sediment less than 350 feet (107 m) thick. These regions are separated by a narrow area of sediment greater than 400 feet (122 m)

thick, trend north-south, and lie to the west of the approximate axis of the Little Knife Anticline.

Subdivisions of the Inyan Kara Formation

Three members of the Inyan Kara were identified and correlated using borehole geophysical logs. Each member was given an informal name.

Member "A"

Identification of member.-- The lowermost member is present in all parts of the study area, and is characterized by sandstone units that are difficult to correlate. As shown in the detailed cross-sections J-J' and K-K' (Plate VII), an individual, well-developed, relatively "clean" sandstone exhibits abrupt lateral change in log character on both the Gamma-Ray log and Dual Laterolog and abruptly pinches out (Interval 5695-5760, Well 103 to Interval 5690-5770, Well 102).

Boundary contacts.-- The basal contact of member "A" is equivalent to the basal contact of the Inyan Kara Formation, which has been described earlier.

The upper contact was chosen at the point where lithologic units become correlatable. The contact is generally abrupt and marked by an abrupt transition from shale (member "A") to a "clean" sandstone or siltstone (member "B") (Well 32, Section H-H', Plate V). In a few cases, the contact is marked by a gradational zone of interbedded siltstone and shale (Well 40, Section E-E', Plate IV), requiring comparison with logs from nearby wells to locate the contact. <u>Thickness</u>.-- The thickness of member "A" varies from less than 200 feet (61 m) in the extreme southern portion of the study area to over 400 feet (122 m) in the west-central portion, and constitutes up to 90% of the total formation thickness. The member thins to the east, north, and southeast, except for an anomalously thin area less than 200 feet (61 m) thick on the northwestern flank of the Little Knife Anticline in the southwestern portion of the study area (T. 144-149 N., R. 98-100 W.).

Linear, oriented bodies of thick and thin sediment dominate the isopach map of the member (Plate X). Orientation of these bodies is east-west in the southern and central portions of the study area and north-south in the northern portion. Bifurcation of relatively thick bodies is towards the north and east.

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Log lithology.-- Using well logs, four basic lithologies were recognized: sandstone, shale, siltstone, and coal.

Sandstone occurs in the study area as units from a few feet (less than 1 m) to greater than 200 feet (61 m) in thickness. On the Gamma-Ray log, four types of sandstone units were recognized:

(1) Upper and lower contacts are sharp with overlying and underlying lithologies; sandstone is relatively homogeneous (Interval 4720-4840 feet, Well 53, Section F-F', Plate IV).

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- (2) Basal contact is sharp, upper contact is gradational with a general decrease in grain size upsection (Interval 5390-5425 feet, Well 36, Section D-D', Plate III).
- (3) Basal contact is gradational; grain size increases upsection with a sharp upper contact (Interval 5130-5200 feet, Well 7, Section A-A', Plate II).

(4) Both upper and lower contacts are gradational (Interval

4620-4710 feet, Well 48, Section F-F', Plate IV). Units of type (1) tend to occur at the base of the member, whereas units of types (2), (3), and (4) tend to occur in the upper portion of "A". The relatively wide separation of the Shallow and Deep Laterolog curves indicates that permeability of these sandstones is relatively high (Fig. 7).

Siltstone occurs in beds less than 5 feet (1.5 m) thick interbedded with shale, either in a gradational zone between sandstone and shale units (Interval 5210-5240, Well 74, Section H-H', Plate V) or in intervals composed dominantly of shale (Interval 4890-4930, Well 81, Section H-H', Plate V).

Shale is present in beds varying from a few feet (less than 1 m) to over 150 feet (46 m) thick. The lower contact is generally gradational with sandstone, whereas the upper contact may either be gradational with or truncated by an overlying sandstone unit. In a few cases, shale is overlain by coal.

Coal occurs in beds less than 10 feet (3 m) thick, and generally overlies sandstone units that coarsen upward.

<u>Core lithology</u>.-- Of the well core in the study area, approximately 55 feet of member "A" was retrieved. Approximately 40 feet (12 m) of core came from Interval 1 and 15 feet (4.5 m) from Interval 2 (Fig. 8).

The basal cored interval of the member includes a generalized vertical sequence of sandstone, siltstone, shale, and coal. The upper cored interval of the member contains shale.

The contact between the Swift Formation and the overlying Inyan Kara is very sharp, with an abrupt transition from black, calcareous, pyritic shale (Swift) (Fig. 9) to a coarse-grained sandstone (Inyan Kara) (Fig. 10).

Sandstone is present in the lower 35 feet (10.6 m) of the member as a sequence of individual units 1-17 feet (0.3-5.2 m) thick. The lower 8 feet (2.4 m) of the sandstone is coarse- to medium-grained, with the exception of a shale-pebble conglomerate 1 foot (0.3 m) above the basal contact of the member (Fig. 11). The upper section of the sandstone decreases in grain size upwards from medium to very fine. The dominant constituent of the sandstone is quartz, which occurs as subangular grains with a dusty gray coating (possibly clay). Other constituents noted, in decreasing abundance, are shale fragments, "white specks" (possibly altered feldspar), and discontinuous laminae of dark, possibly organic, material. The sandstone is moderately friable and does not react to acid, suggesting a weak cement. Sedimentary structures include cross-bedding in the coarse- to medium-grained portion (Fig. 12) and ripple cross-lamination (Fig. 13) and climbing ripples (Fig. 14) in the upper, finer-grained, portion. Except for a few plant fragments in the uppermost portion of the sandstone, no fossils were found.

Figure 9. Shale of the Swift Formation with pyrite crystals (light specks at A). Fig. 8, 4973 feet (1515.8 m).

Figure 10. Contact of the Swift Formation (shale in lower portion of photo) with coarsegrained sandstone of member "A". Fig. 8, 4972 feet (1515.4 m).

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Figure 11. Shale-pebble conglomerate (A) overlying cross-bedded, coarse-grained sandstone in member "A". Fig. 8, 4971 feet (1515.2 m).

Figure 12. Medium-grained, cross-bedded sandstone, with over- and underlying medium-grained, massively bedded sandstone in member "A". Fig. 8, 4957 feet (1510.8 m).



44 Figure 13. Ripple cross-lamination in fine-grained sandstone in the upper portion of member "A". Fig. 8, 4938 feet (1505.1 m). ł Figure 14. Climbing ripples in fine-grained sandstone in the upper portion of member "A". Fig. 8, 4937 feet (1504.7 m).





Figure 15. Ironstone nodules (A) in the uppermost portion of member "A". Note deformation of shale laminae around the nodules. Fig. 8, 4632 feet (1411.8 m).

Figure 16. Lignitic fragment (A) in coarse-grained sandstone with filled pores (B) in member "A". Fig. 8, 4939-4930 feet (1505.6-1502.6 m). ,#¤ ∦:≣

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Two types of shale are found in the member:

- (1) A black, calcareous shale, similar to shale found in the Swift Formation, occurs as sand- to pebble-sized grains or clasts in the lower portion of the sandstone (Fig. 11).
- (2) A medium- to dark-gray, thinly laminated shale is present in both the upper and lower portions of the member (Fig. 14). In the upper portion, both iron-oxide staining and ironstone nodules are present (Fig. 15).

Siltstone occurs as quartzose laminae interbedded with shale in the uppermost part of the core.

The uppermost 9 feet (2.7 m) of Interval 1 are missing, but abundant fragments of lignitic coal are found in the core box, indicating that a coal bed approximately 9 feet (2.7 m) thick was cored but not recovered. Some of the lignitic fragments are imbedded in a coarse-grained sandstone that is partially cemented with gypsum (Fig. 16).

Member "B"

Identification of member .-- The middle member is present in all Farts of the study area, and is characterized by units that are correlatable at distances on the order of several miles.

Boundaries.-- The basal contact of member "B" is generally abrupt, With a moderately thick bed of sandstone or siltstone (10-30 feet (3-9 m) thick) overlying a shale unit of member "A" (Well 6, Section A-A', Plate II). In a few cases the contact is within a gradational zone of interbedded siltstone and shale (Well 20, Section C-C', Plate III).

The upper contact of the member is generally within a gradational $_{zone}$ of siltstone and shale grading upwards into shale (Well 63, Section H-H', Plate V). In a few cases the upper contact is marked by a sharp transition from siltstone or sandstone to shale (Well 43, Section H-H', plate V).

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Thickness.-- Member "B" varies in thickness from less than 20 feet (6 m) in the eastern and southeastern portions of the study area to over 150 feet (46 m) in the east-central portion.

An isopach map of the member (Plate XI) shows two distinct depositional patterns in the study area. In the eastern and southern portions of the study area, the member forms a series of roughly parallel, linear bodies of relatively thick sediment that are oriented northeast-southwest. In the northwestern portion of the study area bodies of relatively thick sediment, which trend either east-west or north-south, are flanked by bodies of thinner sediment that bifurcate towards the east and north. An anomalous area of thin sediment less than 50 feet (15 m) thick occurs in the southwestern portion of the study area west of the axis of the Little Knife Anticline.

Log lithology.-- On logs, the member consists of three sedimentary units, each with a distinctive log character. Between two units, there is a transitional zone that shares log characteristics of both units (Wells 41-43, Section E-E', Plate IV). The presence of these transitional zones distinguishes member "B" from the underlying member "A" and permits moderately good identification and correlation of it in the study area.

Three general types of units were recognized:

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- (1) A basal unit of sandstone or thick siltstone in abrupt contact with member "A" overlain by thinly interbedded siltstone and shale grading upwards into shale (Well 11, Section B-B', Plate II).
- (2) Interbedded siltstone and shale in gradational contact with member "A" grading upwards into shale (Well 65, Section G-G', Plate V).
- (3) Interbedded siltstone and shale in gradational contact with "A", coarsening upward, with a sharp upper contact with "C" (Well 42, Section E-E', Plate IV).

Sandstone units in member "B" are relatively "clean" and less than 30 feet (9 m) thick. An exception is in the southeastern portion of the study area, where sandstone units up to 60 feet (18 m) thick are present (Well 16, Section B-B', Plate II). The basal contact of all sandstone units is generally abrupt whereas the upper contact generally grades upward into interbedded siltstone and shale.

Siltstone occurs in beds that are generally less than 5 feet (1.5 m) thick and are usually interbedded with shale

Shale occurs in beds less than 5 feet (1.5 m) thick and are commonly interbedded with siltstone. Radioactivity of the shales generally increases up section.

<u>Core lithology</u>.-- A complete section of member "B" was retrieved in Interval 2 (Fig. 8). The section consists of interbedded laminae of siltstone and shale. Siltstone laminae vary in thickness from 3-4 inches (7-10 cm) to less than 0.04 inches (1 mm). Laminae are generally homogeneous and appear to be composed almost totally of quartz.

Shale occurs in laminae generally less than 1 inch (2.5 cm) thick. Laminae are gray to black, with layers of abundant plant fragments (Fig. 17).

Contacts between laminae are generally sharp (Fig. 18). Sedimentary structures include rip-up clasts (Fig. 19), micro-dikes of siltstone intruding shale (Fig. 20), convoluted bedding (Fig. 21), burrows (Figs. 19 and 20), and truncated bedding surfaces (Fig. 22). In places, bioturbation was extensive enough to destroy most of the original bedding (Fig. 23).

Member "C"

<u>Identification of member</u>.-- The upper member is shale interbedded with siltstone. The member, on the Laterolog, forms a distinct "sawtooth" pattern of alternating high and low resistivity values (Fig. 8).

<u>Boundaries</u>.-- Both the upper and lower contact of the member are generally gradational. The lower contact, with member "B", occurs within a generally gradational zone of siltstone and shale grading upwards into shale, with the contact chosen at the point where the "sawtooth" pattern of the member begins on the Laterolog. The upper contact is equivalent to the contact of the Inyan Kara and Skull Creek Formations, which has been described previously. Figure 17. Plant fragments (dark spots at A) in shale of member "B". Viewed from horizontal cross-section of core. Fig. 8, 4625 feet (1409.7 m).

Figure 18. Regular and irregular lamination of siltstone and shale in member "B". Fig. 8, 4615 feet (1406.6 m).





Figure 19. Horizontal burrows (A) and rip-up clasts (B) in member "B". Fig. 8, 4629 feet (1410.9 m).

Figure 20. Micro-dike of siltstone intruding shale (A) and vertical burrows (B). Fig. 8, 4615 feet (1406.6 m).

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Figure 21. Convoluted bedding in member "B". Fig. 8, 4628 feet (1410.6 m).

Figure 22. Erosional truncation (along A-B) of beds in member "B". Fig. 8, 4629 feet (1410.9 m). .



Figure 23. Heavily reworked sediments, possibly from burrowing, in member "B". Fig. 8, 4593 feet (1399.9 m).

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Figure 24. Laminated shale with laminae and lenses of siltstone in member "C". Fig. 8, 4590 feet (1399 m).


<u>Thickness</u>.-- Member "C" varies in thickness from approximately 45 feet (14 m) in the southwestern portion of the study area to less than 15 feet (4.5 m) in the eastern portion (Plate XII).

In comparison with members "A" and "B", deposition of "C" was relatively uniform, with a general thinning trend from southwest to the northeast. An anomalous thick area is in the southwestern portion of the study area (T. 148-149 N., R. 94-96 W.).

Log lithology.-- On logs, member "C" forms a distinct, uniform pattern in all parts of the study area. On the Gamma-Ray log, the member is composed of high-radioactivity shales interbedded with thin beds of siltstone less than 5 feet (1.5 m) thick. Radioactivity of the shales is generally highest in the middle of the member and decreases in intensity both up and down section.

On the Laterolog, the member forms a distinct "sawtooth" pattern of alternating high and low resistivity values in response to the interbedded lithologies (Well 1, Section A-A', Plate I). This pattern varies consistently with the thickness of the member, with the number of "teeth" decreasing from west to east (Wells 1 to 9, Section A-A', Plate I).

<u>Core lithology</u>.-- Approximately 2 feet (0.6 m) of member "C" was recovered from the uppermost portion of Interval 2 (Fig. 8). The member is composed almost entirely of thin laminae of black shale, with occasional laminae or lenses of quartzose siltstone (Fig. 24). No fossils or sedimentary structures were recognized.

DISCUSSION

Facies Analysis

Introduction

Most of the interpretations regarding facies and facies relationships in this study were made using borehole geophysical logs. This was necessary due to the limited amount of core available and lack of outcrops in the study area. The interpretations of lithologies from logs were generalized; the interpretations of facies and facies environments were, therefore, also generalized.

Member Facies

Member "A".-- As shown in cross-sections J-J' and K-K', member "A" is composed of laterally discontinuous sandstone units that are difficult to correlate over distances of more than a few miles. The discontinuity of these units suggests that abrupt facies changes are characteristic of the member. Halle (1981) described four Gamma-Ray log patterns, similar to those recognized in member "A", that represented units of nonmarine facies (Fig. 25):

- Distributary channel: Blocky, squared-off pattern with sharp upper and lower contacts (Type 1 in member "A").
- (2) Fluvial channel: Fining-upward, "Christmas tree", pattern (Type2 in member "A").
- (3) Crevasse-splay: Coarsening-upward, "inverted Christmas tree" pattern, often capped by a coal (Type 3 in member "A").

(4) Levee: Sawtooth-shaped pattern (Type 4 in member "A").

Figure 25. Sketches of Gamma-Ray log patterns recognized by Halle (1981) as representing nonmarine facies. ¢

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Pettijohn, Potter, and Siever (1973) compiled a list of physical characteristics of fluvial sandstone units that are similar to those recognized on logs and in well core of member "A":

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- (a) Abundant shale pebbles and shale-pebble conglomerates.
- (b) Carbonaceous debris commonly present.
- (c) Faunal content low to absent.
- (2) Sedimentary Structures
 - (a) Asymmetrical ripple marks.
 - (b) Abundant, well-oriented cross-bedding.
- (3) Size, Shape, and Orientation
 - (a) Commonly elongate, with widths ranging from a few tens of feet to composites of 30 miles.
 - (b) Dendritic, anastomosing, and bifurcating patterns present.
- (4) Associated Lithologic Types
 - (a) Vertical
 - (i) Overlying silty shales, commonly of fluvial origin.Peat or coal commonly present.
 - (ii) Basal contact sharply disconformable.
 - (iii) Multistory sandstone bodies.
 - (b) Lateral
 - (i) Silty shale and siltstone, commonly with abundant carbonaceous material.
 - (ii) Multilateral sandstone bodies, with correlation being generally difficult.

The similarity of characteristics of member "A" and those of nonmarine facies described by Halle and Pettijohn, Potter, and Siever suggests that "A" represents a nonmarine facies of the Inyan Kara Formation. The dominant occurrence of presumed distributary-type deposits at the base of "A" suggests that the lower portion of the member represents a predominately deltaic facies. Log patterns indicating fluvial channel, crevasse-splay, and levee deposits imply that the upper portion of "A" is a fluvial facies.

<u>Member "B"</u>.-- Based upon comparisons in log character the units described in member "B" are similar in shape to the fluvial channel, crevasse-splay, and levee facies of member "A". The units in member "B", however, differ from those in "A" with respect to continuity; the units in "B" are correlatable with relatively high reliability over fairly large distances. The continuity of the units in "B" and the gradational "transition zones" between units suggests that facies in the member are fewer and reflect conditions of more uniform deposition.

Serra and Sulpice (1975) described three Gamma-Ray log patterns, similar to those in member "B", that were interpreted to correspond to units representative of marginal-marine facies (Fig. 26):

- (1) Transgressive sand: Sharp basal contact with a fining-upward pattern (Type 1 of member "B").
- (2) Barrier-bar: Transitional lower contact, coarsening upward, with a sharp upper contact (Type 3 of member "B")
- (3) Tidal channel or tidal flat: Jagged, sawtooth-shaped pattern with gradational upper and lower contacts (Type 2 of member "B").

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Figure 26. Sketches of Gamma-Ray log patterns recognized by Serra and Sulpice (1975) as representing marginal-marine facies.

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Evidence for marginal-marine deposition of member "B" is supplemented by the cored interval of the member. The laminated siltstone and shale units in the core are similar to sediments representative of tidal flat facies described by Reineck and Singh (1980). Although the Gamma-Ray log that accompanies the well core is not of good quality (Fig. 8), the log does resemble the jagged, sawtooth-shaped pattern ascribed to tidal flat or tidal channel facies.

The units recognized in member "B" and the described core of "B" are similar to the trangressive, barrier-bar, and tidal flat or tidal channel facies described by Serra and Sulpice and Reineck and Singh. The similarity of these units implies that member "B" represents a marginal-marine facies of the Inyan Kara Formation.

<u>Member "C"</u>.-- Member "C" differs from members "A" and "B" in that "C" maintains a consistent log character throughout the study area, suggesting that the member represents a single facies. Persistent lateral continuity of units over large areas is generally considered to reflect rocks of a marine facies (Serra and Sulpice 1975). The presence of thin beds of siltstone interbedded with marine shales in the member suggests that periodic influxes of terrigenous material entered the marine environment. The widespread nature and continuity of the siltstone beds in "C" suggest that the siltstone beds were being distributed at a relatively shallow depth, implying that the member was deposited under shallow marine conditions.

Facies Relationships

Due to a lack of outcrops and well core in the study area, no direct evidence of facies relationships between the Swift Formation, members of the Inyan Kara, and the Skull Creek Formation was observed. Indirect evidence for facies relationships between these units is implied, however, by sequences of the interpreted facies of members "A", "B", and "C".

The suggestion of a deltaic facies for portions of the lower part of "A", plus the presence of transitional lower boundary contacts, suggest that the upper portion of the Swift Formation, which has been interpreted to represent a marine facies (Gerhard and others 1982), and the lower portion of member "A" may be facies of a regressive sequence, as illustrated in Fig. 27.

The interpretation of a sequence of fluvial (member "A"), marginalmarine (member "B"), nearshore marine (member "C"), and open marine (lower Skull Creek) facies suggests that the three members and the Skull Creek may be facies of a transgressive sequence, as illustrated in Fig. 28.

Figure 27. Diagrammatic sketch showing suggested conformable relationship between member "A" and the Swift Formation.

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Figure 28. Diagrammatic sketch showing suggested conformable relationships between members "A", "B", "C", and the Swift Formation.



Environments of Deposition

Environments of members "A", "B", and "C"

<u>Member "A"</u>.-- The deltaic and fluvial facies of member "A" suggest that the member was deposited in two separate phases. In the lower portion of the member, where log patterns indicative of distributary-type deposits tended to occur, the deltaic facies implies that the lower portion of "A" was probably deposited in a deltaic environment where distributary channels were dominant. No specific deltaic type was recognized. In the upper portion of the member, where log patterns from fluvial channel, levee, and crevasse-splay deposits tended to occur, the upper portion of "A" was probably deposited in a fluvial environment.

The interpretation of a facies relationship between the deltaic facies of the lower part of member "A" and the marine facies of the Swift Formation implies that the lower portion of "A" was deposited during regression of the epicontinental sea. The relationship between the fluvial facies of the upper portion of member "A" and the marginalmarine facies of member "B" implies that the upper portion of "A" and member "B" are part of a transgressional sequence. No boundary between the regressive and transgressive phases of member "A" was recognized.

Structural and possible paleotopographic influences on the deposition of member "A" in the study area are reflected in an isopach map of the member (Plate X). Two trends of relatively thick sediment, greater than 300 feet (91 m) thick, bifurcate towards the east and

north. The thickness and bifurcation patterns of these areas, along with a general thinning trend of the member to the east, suggest that they may have been major drainage patterns, with sediment transport from the north and east towards the west in the study area. The anomalously thin area on the northwestern flank of the Little Knife Anticline implies that the Little Knife was a slightly positive feature during the deposition of "A", tending to deflect stream flow north and south of the anticline. The Nesson Anticline appears to have not been a significantly positive feature during the deposition of member "A".

<u>Member "B"</u>.-- The interpretation of member "B" is that the member was deposited in a marginal-marine environment. The three recognized facies from well logs (transgressive, tidal flat or tidal channel, and barrier-bar) indicate a lack of deltaic sediments either because the influx of sediments was relatively low or because transgression occurred relatively rapidly. The "transitional zones" noted between the facies suggest that simultaneous deposition of all three facies was occurring during the deposition of "B". The interpreted sequence of nonmarine (member "A"), marginal-marine (member "B"), and shallow marine (member "C") facies implies that "B" is part of a transgressive sequence.

The isopach map of member "B" (Plate XI) reflects the structural and possible paleotopographic controls that influenced the deposition of the member in the study area. In the northwestern portion of the study area, a region of relatively thick sediment, greater than 60 feet (18 m) thick, trends north-south and occurs in the same general area as the

drainage system described for member "A". Log patterns in this area (Wells 29 and 30, Section D-D'; Well 41, Section E-E' and Wells 50, 51, 52, and 53, Section F-F', Plate IV) imply that it was a tidal channel or tidal flat; the shape and orientation of the area imply that it may have been a former river valley. Elsewhere in the study area, log patterns suggest that transgressive sands were dominant in the study area, followed by tidal flat and barrier bar deposits. In the southern portion of the study area the northeast-southwest-trending areas of relatively thick sediment contain log patterns that imply all three facies were present. Possible explanations for the thick accumulation of sediments in these areas is that they either represent a stillstand or a period of high sediment influx which, coupled with subsidence of the Williston Basin, would allow sediments to accumulate. The general thinning trend of the member to the east suggests that transgression of the Cretaceous sea was from west to east in the study area, which is in agreement with previous work (McGookey and others 1972).

Structurally, the interpreted direction of transgression suggests that the Williston Basin was a negative feature during the deposition of member "B". The north-south-trending, anomalously thin area west of the axis of the Little Knife Anticline suggests that the Little Knife was a slightly positive feature during deposition of the member. The Nesson Anticline appears not to have had a significant effect on the deposition of member "B".

<u>Member "C"</u>.-- The interpreted marine facies of member "C" suggests that the member was deposited in an offshore environment. The

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relatively high radioactivity response of the shales in "C" suggests that the shales have a high organic content.

The correlation of a decrease in the number of siltstone beds on the Laterolog with a decrease in thickness suggests that the beds reflect the progression of the depositional environment of "C" across the study area. The widespread extent of the siltstone units in the member suggests that the depositional environment of "C" was shallow and covered a wide area.

The interpreted sequence of marginal-marine (member "B"), shallow marine (member "C"), and deeper marine (Skull Creek) facies, plus the interpreted gradational contacts between member "C" and the Skull Creek Formation, imply that "C" is also part of a transgressional sequence.

The isopach map of member "C" (Plate XII) reflects structural and possible paleotopographic influences on the deposition of the member. For the most part, the deposition of "C" was relatively uniform in the study area. The member thins to the east and northeast, suggesting that transgression in the study area was from west to east. The irregular shape of some of the contours may be the result of paleotopographic influences that are accentuated by the small contour interval. The irregular contours may also be due, in part, to differential compaction in the member. The general thinning trend from west to east suggests that the Williston Basin was a negative feature during the deposition of "C". The anomalously thick area northeast of the Little Knife Anticline in the southwestern portion of the study area implies that the Little Knife continued to be a slightly positive feature during the deposition of member "C". The Nesson Anticline appears to have had no significant effect upon deposition.

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Depositional History of the Inyan Kara

As a general regression of the epicontinental sea progressed across the study area from east to west in Late Jurassic-Early Cretaceous time, deposition of the Swift Formation ceased. Fluvial systems developed, transporting sediment from the east and north into the study area and depositing the sediments of member "A". The relatively high frequency of distributary-type log patterns in the lower portion of the member suggests that, during regression, deltas may have been present.

With the transgression of the epicontinental sea from west to east in the Early Cretaceous in the study area deposition of member "A" was succeeded by that of member "B". Log patterns of transgressive, tidal channel or tidal flat, and barrier-bar deposits suggest that member "B" was deposited in a marginal-marine environment. Temporary halts in the transgression of the Cretaceous sea or increased influx of sediments and subsidence of the Williston Basin are suggested by the thick accumulation of sediments in the southern portion of the study area.

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As further transgression of the Cretaceous sea moved the shoreline eastward, deposition of member "B" ceased and was gradually replaced by member "C". Regular occurrences of siltstone beds in "C" suggest that Periodic influxes of terrigenous material occurred over most of the study area. Further transgression gradually replaced the deposition of "C" with deposition of the Skull Creek.

Structurally, deposition of the members of the Inyan Kara were Ontrolled mainly by the Williston Basin, which was a negative feature hat oriented stream flow for member "A" to the west and controlled ransgression of the Cretaceous sea from west to east during the

deposition of "B" and "C". The only recognized positive structural feature in the study area was the Little Knife Anticline, which caused a slight thinning of each member over or near its axis.

Correlations of the Inyan Kara

Correlations from the study area to the Black Hills

One question that arises in a stratigraphic problem is the lateral extent of a stratigraphic unit. No direct correlation of the Inyan Kara from the Black Hills to the study area was possible due to the lack of available well logs between the Black Hills and the southwestern border of North Dakota. Indirect interpretations were made, however, from comparison of the rocks of the Lakota-Fall River interval in the Black Hills to the members in the study area and from an indirect crosssection from the Black Hills to the study area.

Darton (1901, 1909) first named and described the Lakota-Fuson-Dakota section in the Black Hills. He described the Lakota Formation as a unit of coarse-grained and cross-bedded sandstone with local beds of coal and "partings of shale of no great thickness" lying unconformably upon the "Beulah Shale" (Morrison) or the Unkpapa sandstone. Darton described the Fuson Formation as a nonfossiliferous sequence of fine sand and clay. He described the Dakota Formation as a sandstone with both thin and massive bedding.

Waage (1959) and Robinson and others (1964) provided more detailed descriptions of rocks in the Lakota-Fall River interval. The Lakota Formation (Darton's Lakota and Fuson) was described as sandstone,

siltstone, shale, and coal. Sandstone units in the Lakota were conglomeratic, coarse-, medium-, and fine-grained with cross-bedding, cross-lamination, and massive bedding. The Fall River Formation (Darton's Dakota) was described as a dominantly sandy unit, generally fine-grained, with trails, burrows, casts, and abundant laminated, cross-laminated, and ripple-marked bedding. Above the Fall River, in the lower portion of the Skull Creek Formation, was a sequence of interbedded siltstone and shale.

Rice (1977) constructed cross-sections from the Black Hills to south-central South Dakota and from south-central South Dakota to the study area, providing an indirect cross-section from the Black Hills to the study area. He recognized the Lakota Formation as the lowermost sandstone unit resting unconformably upon Jurassic rocks, the Fall River (Darton's Dakota) as the uppermost sandstone unit lying below the Skull Creek Formation, with the intervening strata designated as the Fuson Formation. Above the Fall River, Rice named and correlated a sequence of interbedded siltstone and shale in the lower Skull Creek Formation as the "basal silt".

Comparisons of stratigraphic position and depositional sequence between the Lakota-"basal silt" interval (of Rice) in the Black Hills and members of the Inyan Kara Formation in the study area suggest that equivalent units may exist in the two areas.

The Lakota Formation of Waage (Darton's and Rice's Lakota-Fuson interval) and member "A" in the study area lie at the bottom of the basal Cretaceous clastic sequence and contain sedimentary features that suggest deposition in a nonmarine environment. The similarity between

these units suggests that the Lakota Formation (of Waage) and member "A" may be approximately equivalent.

The Fall River Formation and member "B" lie above units interpreted to have been deposited in a nonmarine environment (the Lakota Formation and member "A", respectively) and have sedimentary features that suggest the units were deposited in a marginal-marine environment. The similarity betweeen these units suggests that the Fall River Formation and member "B" may be approximately equivalent units.

The "basal silt" (of Rice) and member "C" both lie, in conformable contact, above units interpreted to have been deposited in a marginalmarine environment (the Fall River Formation and member "B", respectively). Both units are composed of interbedded siltstone and shale that suggest deposition in a nearshore marine environment, although Rice has included all siltstone units in the Skull Creek as part of the "basal silt"; member "C" contains only the interbedded siltstone and shale that form the distinctive "sawtooth" pattern on the Laterolog, which occurs in the lower portion of Rice's "basal silt". The similarities between the lower portion of the "basal silt" and member "C" suggest that the units may be approximately equivalent.

Figure 29. Diagrammatic sketch showing the unconformable relationship of member "A" with the Swift Formation as suggested by Anderson and Carlson (1966) and Gerhard and others (1982).

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Analysis of the Swift-Inyan Kara contact

An unconformity in clastic sequences is, by definition, a substantial break in the geologic record that represents an interruption in the depositional sequence (Krumbein and Sloss 1963). Previous work on the Inyan Kara in North Dakota has suggested that the Swift-Inyan Kara contact represents a regional unconformity (Hansen 1955, Rice 1977, Gerhard and others 1982), as illustrated in Fig. 29. In this study, no conclusive evidence was found for a regional unconformity in the study area.

Previous works have cited the pinching out of units in the Swift, such as the one shown in Section C-C', and abrupt changes in lithology from shale (Swift) to sandstone (Inyan Kara) as evidence of an unconformity. In this study, the suggestion of a facies relationship between the Swift and the Inyan Kara offers an alternate interpretation to this cited evidence. The pinching out of units in the Swift, instead of representing erosional truncation, may represent the lateral extent of the units. The abrupt change in lithologies, in which the overlying sandstone units have log patterns that suggest a distributary-channel facies, may not represent an erosional contact; instead, it may represent an abrupt facies change in a marine-deltaic facies sequence.

Evidence for an unconformity in portions of the study area is based upon the observed lithologic changes in well core. The Swift-Inyan Kara contact in the core (Fig. 8) is marked by an abrupt change from shale to coarse-grained sandstone. In the lower portion of the sandstone the presence of sand- to pebble-sized grains of shale, which resemble the lithology of the shales in the Swift, suggests that erosion of the Swift

was occurring nearby, which would imply the presence of a nearby unconformable contact. Unfortunately, the log that accompanies the well core (Fig. 8) is of poor resolution and is not suitable for comparison with other logs.

Evidence for a conformable contact is based on the gradational contacts recognized in the study area (formation contact type 3), which suggests a gradual facies change from marine (Swift) to nonmarine (Inyan Kara), as illustrated in Fig. 25.

CONCLUSIONS

- (1) The Inyan Kara Formation is a clastic unit of sandstone, siltstone, shale, and coal. It varies in thickness from approximately 225 feet
 (69 m) to 500 feet (152 m) in the study area.
- (2) Using well logs, the Inyan Kara can be separated into three members that are correlatable throughout the study area.
- (3) The lowermost member, "A", is a unit of sandstone, siltstone, shale, and coal. It varies in thickness from approximately 200 feet (61 m) to 400 feet (122 m) in the study area.
- (4) Log patterns of sandstone units in the lower portion of "A" suggest that the lower portion of the member was deposited in a dominantly deltaic environment. Log patterns and the discontinuity of sandstone units in the upper portion of "A" suggest that the upper portion of the member was deposited in a dominantly fluvial environment.
- (5) The middle member, "B", is a unit of sandstone, siltstone, and shale. It varies in thickness from approximately 20 feet (6 m) to 150 feet (46 m) in the study area.
- (6) Log patterns in member "B" suggest that the member was deposited in a marginal-marine environment.
- (7) The upper member, "C", is a unit of interbedded siltstone and shale. It varies in thickness from approximately 15 feet (4.5 m) to 45 feet (14 m).
- (8) The presence and lateral continuity of widespread, thin siltstone beds interbedded with marine shale in "C" suggests that the member was deposited in a shallow marine environment.

- (9) Thickness trends in members "A", "B", and "C" suggest that the Williston Basin and the Little Knife Anticline were active structural features during the deposition of the Inyan Kara Formation in the study area.
- (10) Thickness trends in members "A", "B", and "C" suggest that regression of the epicontinental sea in Late Jurassic-Early Cretaceous time was from east to west and that the initial transgression of the Cretaceous sea was from west to east in the study area.
- (11) The upper portion of the Swift Formation and the lower portion of member "A" are suggested to be facies of a regressional sequence.
- (12) The upper portion of member "A", members "B" and "C", and the lower portion of the Skull Creek Formation are suggested to be facies of a transgressional sequence.
- (13) Comparison and indirect correlation of the Lakota-"basal silt" interval (of Rice) in the Black Hills with members "A", "B", and "C" in the study area suggest that the Lakota (or Lakota-Fuson) Formation is approximately equivalent to member "A", the Fall River Formation is approximately equivalent to member "B", and the lower portion of the "basal silt" is approximately equivalent to member "C".
- (14) Analysis of evidence used to suggest a regional unconformity between the Swift and Inyan Kara Formations in the Black Hills suggests that such evidence is ambiguous and, therefore, not conclusive. Some evidence does suggest, however, that locally both conformable and unconformable contacts are present in the study area.



APPENDICES

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APPENDIX A

LEGAL DESCRIPTIONS OF WELLS USED IN CROSS-SECTIONS

Legal descriptions of wells used in cross-sections

Appendix A consists of the legal descriptions of the wells used in cross-sections A-A' to K-K'. Each description consists of two lines. The first line contains, in order, the cross-section reference number, North Dakota Geological Survey well number, well owner or operator, and well name. The second line contains, in order, well location (1/4 1/4 section-township-range) and elevation of the Kelly Bushing.

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- 7361, Amoco Production Company, Stevens Federal C1-29 SW/NE 29-142-100, KB=2757.
- 2. 8947, Amarex Incorporated, Kanski A 5-1 SE/SE 7-142-99, KB=2708.
- 3. 7384, Crystal Exploration and Production Co., Kuntz no. 11-23 NW/NW 23-142-98, KB=2648.
- 4. 8754, Vanderbilt Resources Corp., Rodrmas 1-25 NE/SE 25-142-97, KB=2636.
- 5. 8403, Supron Energy Corp., F.V. Buresch 1 NE/NE 32-142-96, KB=2602.
- 6477, Amoco Production Co., Fisek 1 SE/SW 2-14SE/SW 2-95, KB=2287.
- 7. 5155, Shar-Alan Oil Corp., Mackey 1 SW/SW 13-143-94, KB=2133.
- 7700, Shar-Alan Oil Corp., Walter Benz 1 SW/SW 32-143-93, KB=2221.
- 9. 8115, Keldon Oil Co., Dressler 1 NE/SW 24-142-92, KB=2277.
- 10. 7313, Bello Petroleum Corp., Sheep Creek BN 10-33 NE/SW 33-146-100, KB=2574.
- 11. 7611, Penzoil Exploration and Prod., Grassy Butte 21-21 Federal NE/NW 21-146-99, KB=2629.
- 12. 8535, W.H. Hunt Trust Estate, Brockmier 1 SE/SE 3-146-98, KB=2552.
- 13. 6335, Amoco Production Co., Roquette 1 SW/SE 8-146-97, KB=2509.
- 14. 4611, Helmerich and Payne, Inc., North Dakota State 1 SW/SW 36-146-96, KB=2435.
- 15. 6887, Amoco Production Co., Richardson 1 SW/NE 35-146-95, KB=2325.
- 16. 6448, Smokey Oil Co., O'Neill 11-24 NW/NW 24-146-94, KB=2256.
- 17. 7885, American Natural Gas, F.L.B. Askew 1-14 SE/SE 14-146-93, KB=2406.
- 18. 5532, Cities Service Oil Co., Goetz A-1 NE/NE 33-146-91, KB=2146.

- 19. 8020, Alpha Resources, Inc., Rogness 1-34 SW/NE 34-150-99, KB=2114.
- 20. 7704, Gulf Oil Expl. and Prod., Shafer State 1-23-3B NE/SE 23-150-98, KB=2021.
- 21. 8343, Consolidated Crude Oil, Skjelvik 4-35 NW/NW 35-150-97, KB=2365.
- 22. 7136, Tenneco Oil Co., Lucking 1-27 NW/SE 27-150-96, KB=2329.
- 23. 8083, Texaco, Inc., Mosholder 4 SW/NE 7-150-95, KB=2381.

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- 24. 7673, Helmerich and Payne, Inc., Matthew 1-20 SW/SW 20-150-94, KB=2231.
- 25. 4113, Texaco, Inc.--Skelly Oil Co., Fort Berthold Allottees 1-A SE/NW 4-150-93, KB=2198.
- 26. 7457, Apache Corp., Apache-Grace 1-20 NW/NE 20-150-92, KB=2141.
- 27. 4386, Empire State Oil Co. et al, Vowerk 1 SE/SE 28-151-90, KB=2216.
- 28. 8265, Northwest Exploration Co., 3 Long Creek SW/SE 36-154-99, KB=2252.
- 29. 7915, Brent Exploration, Inc., Seaton 14-30 SE/SW 30-155-98, KB=2278.
- 30. 7931, Mapco, Inc., MCGA 14-33 SE/SW 33-155-97, KB=2124.
- 31. 1266, Amerada Petroleum Co., Woodrow Sven 1 SE/SW 1-154-96, KB=1955.
- 32. 6687, Amerada Hess Corp., State E 32-6-1 SW/NE 16-154-95, KB=2300.
- 33. 3227, Amerada Petroleum Corp. et al, N.D. "N" 1 SE/SE 16-155-94, KB=2030.

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- 34. 2273, Stewart Petroleum Co., Cvancara 1 NW/SW 15-155-93, KB=2360.
- 35. 2816, Davis Oil Co., Len Corkuff 1 SW/SE 12-154-92, KB=2389.
- 36. 6695, Donald C. Slawson, Hill 5-1 NE/NW 5-154-91, KB=2305.

- 37. 8069, Marathon Oil Co., Jensen 12-44 SE/SE 12-154-90, KB=2213.
- 38. 3252, Hunt Oil Co., Annie S. Hoover et al 1 NE/NW 3-158-99, KB=2150.
- 39. 3274, H.L. Hunt, Carl T. Solem 1 NW/SE 2-159-98, KB=2306.
- 40. 8239, Lear Petroleum, Oase 1 SW/SW 17-158-97, KB=2298.

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- 41. 8979, Ranger Oil Co., Herfindahl 11-22 NE/SW 22-158-96, KB=2420.
- 42. 5935, Amerada Heess Corp., NDCA Deep Unit 1 NW/NE 21-158-95, KB=2513.
- 43. 5350, Amerada Hess Corp., TMU L146X NE/SW 7-158-94, KB=2368.
- 44. 4335, I.J. Wilwhite, Calkota, and Assoc., Honrud 1 NW/NE 10-159-93, KB=2290.
- 45. 6047, Apache Corp., Edwards 1-9 SE/SE 9-159-92, KB=2467.
- 46. 5786, Bralorne Intl., Inc., Bralorne et al 15-19 Lumley SW/SE 19-158-91, KB=2379.
- 47. 6677, True Oil Co., Halvorson 43-14 NE/SE 14-157-90, KB=2305.
- 48. 8888, Lear Petroleum, Gordon Hall 1 NW/NE 30-161-98, KB=2117.
- 49. 4394, Texaco, Inc., R.W. Redlin (NCT-1) 1 SW/SW 20-161-97, KB=2157.
- 50. 6798, Shell Oil Co., State Rindel 43-16 NW/SE 16-162-96, KB=2141.
- 51. 7989, Keba Oil and Gas Co., Vande Walle 41X-5 NE/NE 5-162-95, KB=1955.

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- 52. 8139, Bills and St. Clair, State Ahre 1 SE/NW 16-162-94, KB=2016.
- 53. 8831, Texakota, Inc., Ulsrud 2 SW/NE 11-161-93, KB=2297.
- 54. 8364, Texaco Oil and Gas Corp., Hermanson 1 NW/SE 13-162-92, KB=1954.

55.	7859,	Crown	Central	Petroleum,	Arthur	Μ.	Johnson	1-A
	NW/NW	34-16	1-91, KB	=1975.				

- 56. 4599, The Anshutz Corp., Inc., Ormiston 1 SW/SE 25-162-90, KB=1957.
- 57. 3044, Amerada Petroleum Corp., Marie Selle Tract 1-1 NE/NE 27-143-92, KB=2200.
- 58. 8235, Santa Fe Energy Corp., State Coyote Creek 1-36 SE/SE 36-144-92, KB=2258.
- 59. 7978, Terra Resources, Inc., Tozier 1-17 NW/SE 17-145-91, KB=2223.
- 60. 4375, Empire State Oil Co., Youngbear-Sanderson 1 NW/NE 1-149-90, KB=2064.
- 61. 793, Mobil Producing Co., Solomon Birdbear et al 1 SE/SW 22-149-91, KB=2052.
- 62. 7783, Home Petroleum Co., Tribal 1-1 SE/NW 1-150-90, KB=2212.
- 63. 4747, Miami Oil Producers et al, Stolpman 1 NW/NW 30-152-90, KB=1881.
- 64. 6834, Marathon Oil Co., Croft 1 NW/SE 27-155-91, KB=2315.

- 65. 6515, Brownlie, Wallace, Armstrong, and Bander, Jaha 17-11 NW/NW 17-156-91, KB=2340.
- 66. 4433, Union Oil Co. of Calif., Dave Linburg 1 SE/NE 24-159-91, KB=2385.
- 67. 6028, Apache Corp., Apache 1-30 Masters SE/SE 30-160-90, KB=2336.
- 68. 1082, Calvert Drilling, Inc., Jepsen 1 NE/NE 30-161-90, KB=2075.
- 69. 5157, Shar-Alan Oil Corp., Heiser 1 NE/SE 36-143-96, KB=2112.
- 70. 8374, Abode Oil and Gas Corp., Federal Killdeer 41-4 NE/NE 4-144-96, KB=2435.

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- 71. 8431, Amoco Production Co., Sahaydak 1 SW/SW 21-145-96, KB=2573.
- 72. 8077, Mesa Petroleum Co., Pelton 1-10 10-147-96, KB=2417.
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73.	8448, NE/NE	Samedan Oil Corp., Lost Bridge State 1-16 16-148-96, KB=2394.
74.	5936, NW/NE	Ashland Exploration Co., Nelson 1-29 29-149-95, KB=2294.
75.	8631, SE/NW	Texaco, Inc., A.M. Nelson (NCT-1) 2 5-151-95, KB=2386.
76.	8100, SE/NE	Energetics, Inc., Hazen 42-3 3-152-95, KB=2308.
77.	6433, NW/SE	Edward Mike Davis-DBA-Tiger Oil Co., Siguardson 33-23 23-153-95, KB=2344.
78.	6362, SW/SW	Amerada Hess Corp., Marvin Iverson 23-18 18-155-95, KB=2305.
79.	379, 4 SW/SE	Amerada Petroleum Corp., Dena Svor Unit 3 9-156-95, KB=2319.
80.	7832, SW/SW	Dome Petroleum Corp., Neset 1-28 28-157-94, KB=2203.
81.	5411, SW/NE	Hunt Oil Co., NTMU 1-19 6-159-94, KB=2364.
82.	4101, SW/SE	Geochemical Surveys, Smith 1 5-160-94, KB=2379.
83.	6607, NE/NE	North Central Oil Corp., Priebe State 1 5-161-95, KB=2405.
84.	6418, SE/NW	Tenneco Oil Co., B-N 2-29 29-143-100, KB=2647.
85.	6395, NW/SE	Apache Corp., Federal 33-31 31-144-100, KB=2403.
86.	7501, NE/SE	Amoco Production Co., Federal "B" 1 25-145-100, KB=2556.
87.	8341, NE/SW	Tenneco Oil Co., Meinhart USA 1-34 34-147-100, KB=2129.
88.	7700, SW/NE	Texas Oil and Gas Corp., Sperati Federal 1-31 31-148-100, KB=2402.
89.	7943, NE/NW	Amoco Production Co., Hamre 1 23-149-99, KB=2380.
90.	8092, SW/NE	Texaco, Inc., R.T. Lattin 1 27-151-99, KB=2278.

- 91. 2849, Lyda Hunt-Herbert Trusts et al, Henry C. Hystad 1 NE/SW 31-152-99, KB=2316.
- 92. 3406, Hunt Petroleum Corp., Emilia Erickson et al 1 NE/NE 10-156-99, KB=2281.
- 93. 3449, Hunt Petroleum Corp., Chester J. Hamers 1 SE/NW 20-157-98, KB=2213.
- 94. 3491, Hunt Petrolum Corp., Joseph Thvedt 1 NW/SE 13-160-98, KB=2345.

- 95. 8915, Amoco Production Co., Lone Butte Federal Amoco A-1 NE/SW 19-147-97, KB=2113.
- 96. 8414, Gulf Oil Expl. and Prod. Co., Carus Federal 1-30-2C SE/NW 30-147-97, KB=2385.
- 97. 8545, Gulf Expl. and Prod. Co., Mormon Butte Federal 1-25-48 NE/SW 30-147-97, KB=2490.
- 98. 8215, Gulf Oil and Expl. Co., Mormon Butte Federal 1-25-36 SE/SE 25-147-98, KB=2512.
- 99. 8624, Gulf Oil Expl. and Prod. Co., Mormon Butte Federal 4-25-4D SW/SW 25-147-98, KB=2439.
- 100. 7234, Marathon Oil Co., Armour 18-14 SW/SW 18-155-90, KB=2286.
- 101. 7369, Marathon Oil Co., Fenborg 19-32 SW/NE 19-155-90, KB=2281.
- 102. 6974, Brownlie, Wallace, Armstrong, and Bander, State 19-43 NE/SE 19-155-90, KB=2284.
- 103. 8157, BWAB, Inc., Harsrad 29-21 NE/NW 29-155-90, KB=2327.

APPENDIX B

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VALUES OF THICKNESS AND ELEVATION FOR THE

INYAN KARA FORMATION

Values of thickness and elevation for the

Inyan Kara Formation

Appendix B consists of Table 1, which is a list of the wells used in preparing the regional isopach and structure contour maps of the Inyan Kara (Plates VIII and IX). The wells are listed by township (TWN), range (RNG), and section (SEC) and include the elevation of the Kelly Bushing (KB), depth to the top of the Inyan Kara (DKIK) and Swift (DJSW) Formations, elevation of the top of the Inyan Kara (EKIK), thickness of the Inyan Kara (TKIK), and the North Dakota Geological Survey well number (NDGS). The asterisks (*) indicate that the wells were used in preparing the isopach maps of members "A", "B", and "C" (Plates X-XII), a listing of which is included in Appendix C.

TABLE 1

VALUES OF THICKNESS AND ELEVATION FOR THE INYAN KARA FORMATION

TWN	RNG	SEC	*	KB	DKIK	DJSW	EKIK	TKIK	NDGS
129	84	28		2338	3328	3447	-990	119	7930
	85	27		2331	3400	3518	-1069	122	6654
	94	7		2648	4482	4725	-1834	243	8091
	98	30		2695	4562	4852	-1867	290	6050
	100	21		2787	4659	5090	-1872	431	6370
	101	6		2925	4682	5065	-1757	383	5619
	102	2		2857	4590	5018	-1733	428	6074
	103	23		2964	4602	4995	-1638	393	5822
	104	28		3039	4575	4965	-1536	390	6600
	105	13		3149	4445	4800	-1296	355	6176
	106	23		3061	4038	4460	-977	422	6203
130	86	1		1995	3190	3290	-1195	100	4969
	88	23		2206	3450	3612	-1244	162	5118
	91	7		2453	4152	4350	-1699	198	6322
	95	28		2804	4675	4870	-1871	195	7642
	100	17		2867	4865	5290	-1998	425	4545
	102	7		2951	4745	5030	-1794	285	6398
	103	3		3030	4822	5050	-1792	228	5951
	104	15		3179	4845	5180	-1666	335	4143
	107	2		2940	3840	4185	-900	345	37 35
131	85	17		2105	3312	3450	-1207	138	4968
	86	9		2009	3238	3390	-1229	152	4953
	87	31		2451	3710	3875	-1259	165	4935
	88	27		2531	3858	4000	-1327	142	509 7
	98	32		2805	4790	5228	-1985	438	7939
	100	5		2892	4945	5228	-2053	283	5772
	101	19		2932	4908	5268	-1976	360	4449
	102	6		3007	4950	5175	-1943	225	4542
	103	34		3043	4892	5116	-1849	224	5904
	104	23		3260	5020	5362	-1760	342	5838
	105	33		3002	4418	4808	-1416	390	8119
	106	4		2889	4128	4419	-1239	291	8365
132	84	7		2066	3212	3350	-1146	138	4966
	85	29		2257	3475	3575	-1218	100	5113
	86	7		2285	3610	3768	-1325	158	6420
	93	16		2556	4483	4702	-1927	219	8206
	104	15		3167	5000	5340	-1833	340	5888
	106	27		2921	4353	4662	-1432	309	4452
133	83	26		1997	3068	3230	-1071	162	232
	89	19		2437	4115	4328	-1678	213	4701
	90	1		2350	4115	4355	-1765	240	3636
	92	21		2508	4452	4710	-1944	258	6413
	93	26		2517	4575	4792	-2058	217	7075

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$\frac{\text{TWN}}{134}$	$\frac{RNG}{97}$	SEC	*	<u>KB</u> 2669	DKIK 4870	DJSW 5238	<u>EKIK</u>	TKIK 368	$\frac{NDGS}{7453}$
	98	5		2776	5035	5350	-2259	315	5881
	100	23		2879	5085	5435	-2206	350	7016
	101	33		2976	5175	5500	-2199	325	4749
	102	9		2897	5032	5310	-2135	278	5933
	106	15		2872	4420	4820	-1548	400	7787
134	83	26		2145	3369	3535	-1224	166	4967
	90	5		2408	4280	4538	-1872	258	5496
	92	3		2442	4495	4780	-2053	285	8143
	93	22		2367	4507	4832	-2200	265	7231
	95	30		2002	4920	5190	-2324	264	5009
	08	11		2029	5135	5415	-2300	196	1905
	100	22		2142	5202	5300	-2401	202	7800
	105	23		2900	1810	5245	-2321	292	6355
		21		2901	4040	7247	- 255	405	10000
135	82	11		2101	3244	3328	-1143	84	4989
	83	34		2124	3334	3472	-1210	138	3859
	92	12		2524	4585	4930	-2061	345	4984
	08	52		2044	4943	5310	-2399	307	0312 E010
	100	16		2000	5105	5525	-2788	290	1161
	101	10		2776	5210	5520	-2300	310	5929
	103	18		2971	5270	5618	-2299	348	4280
136	81	18		1007	3038	3178	_1131	150	5070
1.50	92	15		2429	4528	4865	-2099	337	5447
	93	35		2548	4739	5020	-2182	290	5783
	96	14		2738	5288	5600	-2550	312	7876
	97	19		2680	5250	5552	-2570	302	6795
	99	7		2784	5418	5745	-2634	327	5245
	100	3		2847	5495	5860	-2648	365	5896
	101	23		2810	52 92	5635	-2482	343	6048
	102	4		2633	5160	5510	-2527	350	6855
	105	22		2694	4879	5232	-2185	353	6319
	106	23		2937	5015	5320	-2078	305	7784
137	82	17		1736	3125	3278	-1389	153	5018
	03	34		2201	3020	3720	-1339	100	3978
	88	20 E		2190	3945	4155	-1/55	210	5011
	00	כ 0		2392	4240	4440 1069	-1900	200	1020 51/10
	92 05	У 22		2 3 2 0	4004	7700 5702	-2110	404 288	5142
	90	22		2688	5320	5710	-2003	285	9299 51世名
	98	2		2749	5542	5936	-2793	394	5482
	99	14		2749	5550	5940	-2801	390	6438
	100	2		2800	5515	5822	-2715	307	5195
	102	20		27 19	5382	5742	-2663	360	2357
	103	28		2728	5220	5572	-2492	352	7842

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TWN 127	RNG	SEC	*	<u>KB</u>	DKIK	DJSW	EKIK	TKIK	NDGS
121	104	22		2090	5058	5182	-2026	121 121	6272
	100	26		5052	5050	J-10L	-2020	767	0212
138	82	21		1896	3178	3451	-1282	273	5019
	83	5		1980	3364	3690	-1384	326	5379
	85	6		2076	3755	4075	-1679	320	7770
	86	19		1965	3780	3963	-1815	183	7937
	92	9		2407	4698	5072	-2291	374	6476
	94	23		2554	5062	5405	-2508	343	5203
	97	18		2726	5800	6097	-3074	297	2496
	90	20		2150	5095	6200	-3139	313	6207
	100	29		2090	5560	5895	-2718	335	6601
	101	36		2888	5580	5975	-2692	395	6485
	102	1		2737	5370	5742	-2633	372	5243
	103	29		2765	5239	5600	-2474	361	4490
	104	27		2725	5262	5169	-2537	357	5832
	105	9		2867	5240	5656	-2373	416	4130
139	82	11		1861	3225	3338	-1364	113	2185
	86	30		2204	4038	4234	-1834	196	133
	90	27		2426	4525	4860	-2099	335	1620
	92	16		2494	4830	5180	-2336	350	6797
	93	11		2450	4888	5173	-2432	285	4446 5751
	94	21		2450	5242	5510	-2192	320 250	2/24 7076
	90	20		2530	5548	5802	-2039	254	5464
	97	- 8		2496	5275	5572	-2779	297	6447
	98	15		2544	5418	5785	-2874	367	6369
	99	9		262 9	5614	5972	-2985	358	5014
	100	6		2805	5636	6010	-2831	374	5342
	101	1		2753	5568	5920	-2815	352	5869
	102	11		2523	5210	5562	-2687	352	3573
	103	17		2898	5500	5009	-2602	509	4999
	104	ני ע		2808	5368	5770	-25(0	454	4300
	106	10		2836	5218	5680	-2382	462	6861
120	86	26		22011	20E8	11211	_175#	25.2	5017
140	88	26		2230	4270	4588	-2040	318	7340
	89	5		2284	4512	4888	-2228	368	7818
	93	11		2293	4908	5234	-2615	326	5415
	94	27		2395	5128	5486	-2733	358	6691
	95	13		2474	5400	5738	-2926	338	5361
	96	14		2456	5631	5941	-3085	310	5526
	97 08	- 9 10		2590	5530	5095	-2940	359	5440 5685
	90	21		2512	5582	5881	-2034	202	6820
	100	22		2767	5745	6068	-2914	302	4497
	102	19		2413	5138	5546	-2725	408	6106

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TWN	RNG	<u>SEC</u>	*	KB	DKIK	DJSW	<u>EKIK</u>	TKIK	NDGS
140	103	10		2662	5365	5874	-2703	509	6331
	105	33		2814	5378	5763	-2564	385	5738
141	82	15		1973	3435	3695	- 1462	260	8144
	85	34		2173	3980	4160	-1807	180	6894
	91	8		2113	4715	5040	-2602	325	824
	93	17		2227	5060	5335	-2833	275	5470
	94	24		2255	5156	5464	-2901	308	5274
	95	18		2595	5443	5702	-2848	259	6530
	90	30		2591	5404	5852	-2873	300	5399
	91	l R		2090	5570	5920	-29/2	350	6510
	90	16		2628	5628	5075	-2909	200	6788
	100	26		2556	5429	5790	-2873	361	5861
	101	22		2333	5181	5645	-2848	464	6659
	102	9		2470	5290	5730	-2820	440	7527
	103	31		2726	5485	5985	-2759	500	8460
	104	7		2759	5435	5812	-2676	377	7753
	105	27		2710	5532	5750	-2622	418	5438
142	84	7		2138	3996	4162	-1858	166	4941
	85	14		2193	4018	4224	- 1825	206	3227
	91	6		2000	4760	5076	-2760	316	4727
	92	24	*	2155	4888	5213	-2733	325	5131
	93	1	ж ж	2001	5080	5390	-3019	318	9096
	90	21	¥	2618	5600	5411	-2031	293	7260
	97	23	¥	2583	5546	5870	-2963	324	5621
	98	36	¥	2676	5632	5941	-2956	309	6140
	99	3	¥	2728	5754	6078	-3026	324	7618
	100	22	¥	2720	5591	5869	-2871	278	6914
	101	13		2535	5360	5755	-2825	395	7452
	102	7		2414	5280	5722	-2866	442	5842
	103	3		2595	5845	5905	-2890	420	7255
	104	22		2595	5380	5005	-2705	425	4400
	105	52		2092	2312	5745	-2020	433	1909
143	83	10		1951	3605	3745	-1654	140	4938
	04 86	33		2139	3955	4125	-1816	170	4939
	00	16	¥	2221	4200	4430 5220	-2001	275	9942
	92	23	*	2248	5022	5324	-2774	326	1787
	94	13	¥	2133	5000	5355	-2867	355	5155
	95	31	¥	2214	5221	5567	-2998	346	8275
	96	36	¥	2211	5218	5570	-3007	352	5157
	97	17	¥	2575.	5569	5852	-3015	283	8202
	98	28	¥	2649	5718	6045	-3019	327	6647
	99	2	¥	2722	5760	6090	-3038	330	7348

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<u>TWN</u> 143	RNG 100 101 102 103 105	SEC 29 4 1 25 17	* *	<u>KB</u> 2647 2428 2285 2579 2669	DKIK 5560 5308 5198 5445 5509	DJSW 5906 5670 5552 5882 5885	EKIK -2913 -2880 -2913 -2866 -2840	<u>TKIK</u> 346 362 354 437 376	NDGS 6418 6724 5711 6821 4735
144	85 88 90 91 92 93 95 97 97 99 90 100 102 103 104 105	34 10 11 29 16 7 4 28 30 6 31 4 30 6 14 30 6 14 18 123	* * * * * * * *	1958 2059 1968 2080 2121 2240 2121 2343 2310 2508 2532 2750 2597 2198 2341 2531 2559 2597 2198	3888 4385 4395 4812 4946 5073 5122 5377 5583 5575 5578 5578 5578 5578 5578	4110 4795 4730 5197 5160 5317 5418 5864 5767 6010 5935 6078 5881 5479 5660 5884 5815 5868	-1930 -2326 -2427 -2732 -2689 -2685 -2833 -3001 -3082 -3067 -3075 -3043 -3050 -2981 -2947 -2951 -2973 -2962 -2886	222 410 335 385 350 371 345 358 439 390 427 360 278 303 334 368 380 402 413	4823 377 4808 7616 5233 7346 4748 8153 6489 7426 6214 4035 6310 4419 6947 6508 6562
145	80 81 82 83 85 86 80 91 99 99 99 99 99 99 99 90 101 102	21 8 10 28 29 17 31 5 17 34 17 28 27 1 18 32 7 118 13 11 7 29 17	* * * * * * * * * * *	1944 2013 1875 2008 1699 1948 1924 2203 2141 2203 2141 2221 2327 2573 2577 2614 2680 2346 2118 2432	3344 3515 3478 3748 3560 4020 4162 4898 4914 5156 5427 55727 55727 55727 55729 5518 55318 55318 5538	3560 3758 3995 3846 4380 4582 5148 52249 5180 5555 5730 5555 5792 6118 6095 6250 5761 5520 5948	-1400 -1502 -1603 -1740 -1861 -2048 -2166 -2238 -2459 -2813 -2691 -2706 -2431 -3063 -3154 -3076 -3155 -3135 -2972 -2956 -3106	216 243 268 2280 360 486 335 339 350 416 435 399 201 416 435 446 410	4814 4815 4810 4824 4795 4795 4795 4177 6674 7978 4167 4482 60348 5166 6348 6230 26558 7746

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<u>TWN</u> 145	RNG 103	<u>SEC</u> 19	*	<u>KB</u> 2590	DKIK 5654	DJSW 6120	<u>EKIK</u> -3064	<u>TKIK</u> 466	NDGS 6014
	104 105	12 24		2558 2379	5625 5330	6038 5804	-3067 -2951	413 474	6655 3645
146	81 82 90 91 93 94 95 96 97 98 99 100 101 102 103 104	34 32 35 314 25 314 21 35 8 311 21 15 10 5 105	* * * * * * * * *	1932 2022 2186 2309 2146 2396 2325 2658 2509 2262 2628 2590 2494 2210 2291 2391 2280	3412 3645 4325 5052 4962 5260 5284 5356 5713 5368 5791 5368 5791 5368 5372 5365 5372 5365 5438 5290	3645 3910 4663 5418 5335 5645 5743 6117 5986 6169 6072 5950 5760 5832 5870 5730	-1480 -1623 -2139 -2743 -2816 -2864 -3098 -3055 -3055 -3055 -3092 -3106 -3163 -3128 -3066 -3162 -3074 -3047 -3010	233 265 338 366 373 361 430 404 385 298 378 354 354 390 388 467 432 440	4813 1516 4797 3492 5532 7885 6182 6887 7477 6335 6324 7611 6786 7495 6716 7549 6382
147	80 92 95 96 97 98 99 100 101 102 103 104	11 10 8 23 22 18 36 27 34 21 27 17 29 1	* * * * * *	1880 2048 2297 2379 2525 2055 2513 2587 2129 2196 2222 2216 2327 2090	3203 4903 5196 5510 5703 5120 5586 5729 5274 5380 5382 5356 5401 5168	3452 5312 5910 6066 5452 5940 5568 5684 5765 5770 5824 5690	-1323 -2855 -2899 -3131 -3078 -3075 -3073 -3142 -3146 -3140 -3140 -3074 -3074	249 409 416 400 363 352 354 351 294 304 383 414 423 522	5261 7745 2848 7576 6623 8486 6345 7636 8341 6718 7850 6443 6848 6456
148	89 90 95 96 97 98 100 102 103 104	30 1 18 20 36 9 4 24 7 2	* * * * *	2058 2064 2441 2095 2443 2206 2311 2433 2474 2321 2081	4765 4773 5478 5072 5552 5327 5594 5675 5700 5512 5202	5105 5160 5848 5428 5900 5675 5924 6110 6120 6007 5690	-2707 -2728 -3039 -2977 -3109 -3121 -3283 -3242 -3242 -3226 -3191 -3121	340 387 370 356 348 348 330 435 420 495 488	5352 4375 2352 7614 3178 6597 6524 7580 7648 6387 6383

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TABLE 1--CONTINUED

<u>TWN</u> 149	RNG 86 90 91 92 93 95 96 97 98 99 100 102	SEC 16 5 22 33 24 16 24 2 11 22 35 8 8	* * * * * * * * * *	KB 2227 1989 2102 2363 2149 2294 2499 2212 2232 2127 2207 2397 2212	DKIK 4376 4800 4942 5187 5210 5372 5253 5538 5495 5528 5658 5658	DJSW 4769 5168 5322 5595 5422 5576 5750 5625 5828 5868 5870 6030 5895	EKIK -2149 -2811 -2840 -2824 -2898 -2916 -2873 -3041 -3306 -3368 -3321 -3261 -3276	TKIK 393 368 380 408 375 366 378 372 290 373 342 372 <th>NDGS 5126 1194 793 7214 607 5936 6368 4085 3157 6493 7879 7478 76493</th>	NDGS 5126 1194 793 7214 607 5936 6368 4085 3157 6493 7879 7478 76493
150	80 86 88 90 92 93 94 95 97 98 97 98 97 100 101 102 103	14 27 29 12 1 20 4 33 20 33 20 31 20 31 325	* * * * * * *	2006 2162 1999 2019 2190 2143 2198 2334 2342 2327 2244 2022 2141 2335 2308 2262 2127 2170	3330 4282 4545 4550 4981 5000 5147 5163 5152 5147 5293 5522 5675 5622 5580 5477 5384	3604 4662 4916 4947 5360 5357 5522 5560 5484 5532 5571 5708 5960 6115 6087 5998 5915 5866	-1324 -2120 -2546 -2531 -2791 -2857 -2949 -2829 -2829 -2829 -2758 -2825 -2903 -3271 -3381 -3340 -3314 -3318 -3350 -3214	274 380 371 397 379 357 375 397 384 475 438 4475 438 440 465 438 482	3089 8310 5826 6766 7783 7457 4113 3731 6608 6849 7743 7704 7873 5774 7796 7650 7422 5655
151	81 86 89 90 94 95 96 97 99 101 102 103	19 23 20 34 10 13 34 11 27 23 1 8	* * * * *	2146 2150 2147 2139 2115 1956 2107 2432 2291 2278 2048 2086 2200	3642 4221 4655 4761 4744 4734 5081 5093 5629 5375 5385 5508	3972 4648 5020 5171 5152 5130 5457 5478 6055 5785 5895 5910	-1478 -2071 -2508 -2546 -2646 -2788 -2627 -2649 -2802 -3351 -3327 -3299 -3308	348 427 365 405 408 396 385 426 410 510	5401 5096 5731 4392 4594 3056 4095 7008 4095 7008 4723 7142 1642

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152 02 33 2006 3605 4005 -1519 400 588 83 13 2110 3673 4092 -1563 419 3080 85 7 2122 4170 4525 -2048 355 7554 86 28 2120 4285 4635 -2165 350 5105 87 28 2092 4336 4710 -2244 374 4747 93 16 2020 5004 5427 -2984 424 4061 94 7 2190 4805 5190 -2613 387 3387 95 7 2363 4865 5249 -2502 384 4013 96 14 2451 4957 5348 -2506 391 6323 99 31 2316 5648 6063 -3332 415 2849 101 5 2198 5492 5960 -3294 468 6946 102 35 2276 5575 <	1WN	RNG	SEC	—	KB	DKIK	DJSW	EKIK	TKIK	NDGS
83 13 2110 3673 4092 -1563 419 3080 85 7 2122 4170 4525 -2048 355 7554 86 28 2120 4285 4635 -2165 350 5105 87 28 2092 4336 4710 -2244 374 5313 88 19 2086 4538 4903 -2452 365 2779 90 30 $*$ 1881 4568 4942 -2687 374 4747 93 16 $*$ 2020 5004 5427 -2984 424 4061 94 7 $*$ 2190 4803 5190 -2613 387 3387 95 7 $*$ 2363 4865 5249 -2502 384 4013 96 14 $*$ 2451 4957 5348 -2506 3916 6323 99 31 $*$ 2316 5648 6063 -3332 415 2849 101 5 2198 5492 5960 -3294 468 6946 102 35 2276 5575 6085 -3299 510 6799 104 2 2169 5435 5969 -3266 534 8178 153 80 14 1636 2858 3158 -1222 300 5344 82 15 1795 3286 3658 <t< td=""><td>152</td><td>02</td><td>33</td><td></td><td>2000</td><td>3605</td><td>4005</td><td>-1519</td><td>400</td><td>588</td></t<>	152	02	33		2000	3605	4005	-1519	400	588
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$\begin{array}{cccccccccccccccccccccccccccccccccccc$		104	25		2172	5440	5972	-3268	532	7632
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	154	81	10		1566	2070	3250	_ 1202	280	3237
83 8 1845 3511 3918 -1666 407 2930 84 32 2132 4010 4420 -1878 410 6653 85 6 2219 4245 4460 -2026 215 5389 86 22 2121 4253 4475 -2132 222 4223 88 5 2132 4455 4792 -2323 337 7238 89 29 2162 4705 5070 -2543 365 7918 90 22 2235 4823 5221 -2588 398 7414 91 5 2305 5106 5484 -2801 378 6695 92 15 5232 5632 2632 409 2916	1.2.1	82	31		1797	2378	3664	-1581	286	1891
84 32 2132 4010 4420 -1878 410 6653 85 6 2219 4245 4460 -2026 215 5389 86 22 2121 4253 4475 -2132 222 4223 88 5 2132 4455 4792 -2323 337 7238 89 29 2162 4705 5070 -2543 365 7918 90 22 2235 4823 5221 -2588 398 7414 91 5 2305 5106 5484 -2801 378 6695 92 12 #239 5223 5630 2933 100 2916		83	8		1845	3511	3918	-1666	407	2930
85 6 2219 4245 4460 -2026 215 5389 86 22 2121 4253 4475 -2132 222 4223 88 5 2132 4455 4792 -2323 337 7238 89 29 2162 4705 5070 -2543 365 7918 90 22 * 2235 4823 5221 -2588 398 7414 91 5 * 2305 5106 5484 -2801 378 6695 92 12 * 2328 5623 2623 409 2916		84	32		2132	4010	4420	-1878	410	6653
86 22 2121 4253 4475 -2132 222 4223 88 5 2132 4455 4792 -2323 337 7238 89 29 2162 4705 5070 -2543 365 7918 90 22 * 2235 4823 5221 -2588 398 7414 91 5 * 2305 5106 5484 -2801 378 6695 92 12 * 2329 5232 5630 2933 409 2916		85	6		2219	4245	4460	-2026	215	5389
88 5 2132 4455 4792 -2323 337 7238 89 29 2162 4705 5070 -2543 365 7918 90 22 * 2235 4823 5221 -2588 398 7414 91 5 * 2305 5106 5484 -2801 378 6695 92 12 * 2323 5620 2823 5620 2823 108 2816		86	22		2121	4253	4475	-2132	222	4223
89 29 2162 4705 5070 -2543 365 7918 90 22 * 2235 4823 5221 -2588 398 7414 91 5 * 2305 5106 5484 -2801 378 6695		88	5		2132	4455	4792	-2323	337	7238
90 22 * 2235 4823 5221 -2588 398 7414 91 5 * 2305 5106 5484 -2801 378 6695 92 12 * 2380 5222 5620 2822 408 2816		89	29		2162	4705	5070	-2543	365	7918
91 5 * 2305 5106 5484 -2801 378 6695		90	22	¥	2235	4823	5221	-2588	398	7414
		91	5	¥	2305	5106	5484	-2801	378	6695
32 12 ° 2303 7222 3030 -2033 400 2010		92	12	¥	2389	5222	5630	-2833	408	2816
95 18 * 2052 4297 4750 -2245 453 6184		95	18	¥	2052	. 4297	4750	-2245	453	6184
96 1 * 1955 4240 4702 -2285 462 1266		96	1	¥	1955	4240	4702	-2285	462	1266
99 36 * 2256 5630 6125 - 3374 495 7074		99	36	¥	2256	5630	6125	-3374	495	7074

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TWN	RNG	SEC	*	KB	DKIK	DJSW	EKIK	TKIK	NDGS	
124	100	20		1000	5200	5/00	-3320	492	24/0	
	102	9 h		2100	5160	5900	-3257	517	6002	
	105	4		2309	5459	2910	-2124	517	5992	
155	81	21		1593	2902	3230	-1309	328	5531	
	82	25		1659	3015	3380	-1356	365	6100	
	84	8		1884	3650	3932	-1766	282	6725	
	85	32		2206	4160	4428	-1954	26 8	6610	
	86	16		2234	4305	4618	-2071	313	5468	
	87	15		2219	4362	4671	-2143	309	7612	
	88	35		2135	4414	4776	-2279	362	6179	
	89	19		2034	4501	4861	-2467	360	7368	
	90	21	¥	2317	4835	4854	-2518	377	7220	
	91	27	¥	2315	5032	5450	-27 17	418	6834	
	93	15	¥	2360	5313	5732	-2953	419	2273	
	94	16	¥	2030	4739	5130	-2709	391	3227	
	95	18	¥	2305	4524	4972	-2219	448	6362	
	96	35		2032	4222	4680	-2190	458	6388	
	97	14	*	2253	4938	5362	-2685	424	7089	
	98	21	*	2248	5502	5900	-3254	398	7712	
	99	8	*	2116	5365	5826	-3249	461	7405	
	100	30		1898	5175	5618	-3277	443	7004	
	101	0		2192	5425	5938	-3233	513	6806	
	102	24		2192	5402	5935	-3270	473	7692	
156	80	8		1526	2594	2904	-1068	310	4112	
	81	8		1580	2848	3160	-1268	312	5563	
	82	5		1638	2980	3335	-1342	355	5360	
	83	22		1678	3199	3550	-1521	351	4805	
	84	10		1730	3295	3715	-1565	420	4400	
	85	19		1982	3910	4220	-1928	310	3984	
	86	30		2206	4357	4625	-2151	268	7695	
	87	12		2146	4293	4583	-2147	290	7042	
	90	31	*	2280	4865	5253	-2585	388	7847	
	91	- 17	*	2340	4984	5397	-2544	413	6515	
	92	14	π #	2322	5021	5417	-2699	390	592	
	93	20	×	2310	5230	5040	-2002	402	2333	
	94	20	×	2331	5000	5351	-2009	351	((4)	
	95	2	т ¥	2204	4020 NERE	5009	-2304	441	(099 6107	
	90 07	30 27	¥	2333	5015	2013	-2230	430	2461 5/17	
	71	2 { 10	¥	2281	5045	5000	-2113	100	3106	
	101	16		2168	5310	5772	-3172	1122 1122	3225	
	102	14		2151	5246	5717	-3095	471	7054	
	103	22		2380	5405	5870	-3016	465	6789	
	104	12		2437	5370	5859	-2933	489	7356	

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TWN	RNG	<u>SEC</u>		KB	DKIK	DJSW	<u>EKIK</u>	TKIK	NDGS
157	79	12		1487	2306	2636	-819	330	5078
	80	2		1506	2466	2795	-960	329	5982
	81	24		1535	2640	2961	-1105	321	4409
	82	36		1622	2912	3308	-1290	396	6486
	83	1		1631	2961	3280	-1330	319	5539
	84	2		1770	3312	3692	-1542	380	6502
	85	24		1829	3550	3962	-1721	412	6032
	86	3		1887	3754	4170	-1867	416	5809
	87	1		2053	4155	4470	-2102	315	6978
	88	12		2301	4435	4755	-2134	320	3673
	89	20		2222	4550	4945	-2328	395	4194
	90	19		2384	4930	5305	-2546	375	1406
	91	4	¥	2372	4946	5340	-2574	394	4682
	94	10		2290	4815	5203	- 2525	388	1270
	95	33	¥	2271	4580	5022	-2309	442	6111
	96	30		2324	4955	5357	-2631	394	8296
	97	2	¥	2296	5111	5506	-2815	395	7789
	98	20	¥	2213	5323	5717	-3110	394	3449
	101	17		2191	5180	5702	-2989	522	7834
	103	10		2358	5512	5742	-2854	530	5271
1-0	70	-		1100		0001	0 - 4		
150	79	1		1480	2354	2694	-874	340	5076
	0U 81	12		1511	2543	2040	-1032	297	7609
	82	12		1510	2002	2019	-1000	211	0300
	82	28		1672	2090	2042	1260	302 21/5	1219
	81	19		1755	2110	2201	-1309	345	6100
	85	22		1810	25 27	2810	-1000	3/4	0409 5585
	86	21		1075	2068	1078	-1/10	210	5505
	87	21		1760	2820	9210	-1993	310	60/1
	88	24		2250	1185	4195	-2051	222	25/10
	80	0		2273	4620	5022	-2220	202	3005
	01	10	¥	2370	5017	5270	2628	252	5786
	92	Q		2288	1012	5300	-2654	358	1315
	01	7	¥	2368	1912	5283	-2520	305	5350
	95	21	¥	2513	4000	5320	-2306	420	5935
	96	26	¥	2354	4882	5245	-2528	363	7565
	97	17	¥	2276	5126	5533	-2850	407	8239
	99	.,	¥	2150	4942	5328	-2792	386	3252
	100	1		2133	4920	5375	-2787	455	8180
	101	10		2087	4896	5355	-2809	459	7164
	103	6		2035	4642	5022	-2607	380	5054
					• • • •		.		· · · ·
15 9	79	29		1457	2374	2648	-917	274	b700
	0U 21	29		1495	2445	2780	- 950	335	1004
	20	33		1521	2020	2934	-1105	300	0319
	02 80	34		1501	2150	31/0	-1103	420	2052
	03	33		1040	2920	3201	-1202	349	1051

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TWN	RNG	SEC	<u>*</u>	KB	DKIK	DJSW	EKIK	TKIK	NDGS
159	84	8		1697	3146	3496	-1449	350	6328
	85	31		1828	3566	3882	-1738	316	5615
	86	26		1863	3688	4030	-1825	342	5851
	87	3		1921	3872	4188	-1951	316	2134
	88	32		2141	4344	4663	-2203	319	1843
	90	17		2398	4785	5165	-2387	380	3604
	91	24	¥	2381	4814	5171	-2433	357	4433
	92	9	¥	2467	4910	5282	-2443	372	6047
	- 93	10	¥	2290	4808	5200	-2518	392	4335
	94	6	¥	2364	4750	5147	-2386	397	5411
	95	30	*	2370	4862	5236	-2492	374	3126
	96	12	*	2317	4813	5173	-2496	360	3392
	97	10	*	2301	4984	5377	-2683	393	8039
	98	2	*	2306	4963	5308	-2657	345	3274
	99	2	*	2273	4938	5312	-2665	374	6017
	101	2		2160	4760	5418	-2600	388	5509
	102	15		2295	4895	5330	-2600	435	6162
	103	24		2206	4786	5110	-2580	324	4663
160	79	16		1464	2224	2518	-760	294	4915
	80	19		1511	2402	2732	-891	330	2596
	81	5		1516	2472	2800	-956	328	4192
	82	5		1549	25 85	3000	-1036	415	4431
	83	27		1600	2797	3128	-1197	331	4997
	84	21		1687	3028	3400	-1341	372	6001
	85	27		1743	3325	3653	-1582	328	6446
	86	4		1821	3535	3840	-1714	305	7716
	87	16		1893	3768	4105	-1875	337	6482
	88	36		1952	3962	4363	-2010	401	4097
	89	19		2062	4173	4495	-2111	322	1005
	90	30	*	2336	4631	4978	-2295	347	6028
	91	5	*	2414	4658	4999	-2244	341	5721
	92	21	*	2358	4750	5132	-2392	382	7917
	93	23	*	2371	4805	4895	-2434	359	2684
	94	5	*	2379	4816	5198	-2437	382	4101
	95	3	*	2373	4662	5078	-2289	416	5192
	96	35	*	2290	4745	5083	-2455	338	5009
	97	19	*	2307	4865	5228	-2558	363	5547
	98	10	*	2243	4722	5040	-2479	318	5248
	99	24	*	2236	4775	5070	-2539	295	7110
	100	12		2212	4583	4978	-2471	395	4837
	101	32		2257	4792	5168	-2535	376	3374
	102	26		2263	4735	5148	-2472	413	4962
	103	33		2006	4430	4869	-2430	433	7658
161	79	1		1470	2140	2392	-670	252	6259
	80	9		1500	2312	2512	-812	200	5943
	81	10		1514	2450	2740	-9 36	290	5358

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TWN	RNG	SEC	*	<u>КВ</u>	DKIK	DJSW	EKIK	TKIK	NDGS
161	82	33		1561	2648	3058	-1087	410	4918
	83	17		1611	2785	3168	-1174	383	5455
	84	10		1666	2955	3355	-1285	404	3987
	85	14		1754	3190	3538	-1436	348	6166
	86	22		1829	3486	3786	- 1657	300	6507
	87	4		18 81	3588	3960	-1707	372	5796
	89	32		1994	4075	4400	-2081	325	1842
	90	15	¥	1983	4068	4373	-2085	305	5850
	91	2	¥	1973	4084	4395	-2111	311	4958
	92	15	¥	2422	4658	5060	-2236	402	4317
	93	2	¥	2244	4515	4886	-2271	371	7575
	94	5	¥	2364	4612	4990	-2248	378	5246
	96	25	¥	2364	4720	5065	- 2356	345	2996
	97	20	¥	2157	4603	4937	-2446	334	4394
	98	13	¥	2060	4462	4816	-2402	354	2722
	100	15		2242	4587	4892	-2345	305	2721
	101	3		2244	4542	4957	-2298	415	6751
	102	19		2102	4385	4805	-2283	420	4487
	103	25		2075	4330	4770	-2255	440	6900
162	79	35		1470	2125	2362	-655	237	7270
	80	11		1496	2210	2530	-714	320	5038
	01	32		1520	2415	2/00	-095	345	3034
	82	20		15 30	2440	2100	-902	350	4720
	81	52 1 h		1626	2750	2125	-1147	325	1199
	85	2		1620	2000	3280	-1266	215	5770
	86	20		1837	2905	3765	-1608	320	5779
	87	1		1716	3246	3590	-1530	3111	6501
	89	14		1923	3770	4105	-1847	225	3365
	90	32	¥	1964	3996	4303	-2032	307	7664
	91	34	¥	1975	4103	4410	-2128	307	7859
	92	15	¥	1948	4039	4413	-2091	374	6081
	93	36	¥	2096	4364	4726	-2268	362	6460
	94	21		2050	4302	4675	-2252	373	7528
	95	20	¥	2136	4303	4668	-2167	365	4074
	96	16	¥	2141	4388	4728	-2247	340	6798
	98	5	¥	2221	4482	4853	-2261	371	1900
	100	13		2349	459 3	4950	-2244	357	6541
	101	27		2201	4535	4920	-2334	390	6652
	102	3		2244	4373	4621	-2129	255	3441
	103	26		2145	4330	4730	-2185	400	6429
167		0		1100	2065	0005	- 0 -	200	
103	(Y 80	0 26		1400	2005	2395	-505 670	330	6126
	81	22		1515	2220	2575	-013	255	7208
	82	24		15 20	2224	2684	_801	350	7117
	83	23		1594	2578	2900	_984	322	7176
		-0			-210	-,	501	J	1

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TWN	RNG	SEC	¥	KB	DKIK	DJSW	EKIK	TKIK	NDGS
163	-84	27		1620	2700	3072	-1080	372	7491
	85	30		1736	2989	3393	-1253	404	5649
	86	11		1752	3060	3500	-1308	440	7473
	87	3		1645	3085	3480	-1440	395	6749
	89	23		1907	3565	3850	-1658	315	6877
	90	7		1943	3742	4092	-1799	350	7540
	91	7		1957	3820	4200	-1863	38 0	5950
	92	7		1941	3927	4235	-1986	345	7837
	93	15		1933	4008	4410	-2075	402	5782
	94	6		1921	3967	4323	-2046	356	7773
	95	29		1926	4055	4420	-2129	365	4691
	96	32		1932	4195	4484	-2193	359	7515
	97	26		1957	4145	4516	-2188	371	3903
	98	30		2119	4469	4785	-2350	316	6061
	99	23		2209	4400	4705	-2191	305	5404
	100	27		2170	4310	4630	-2140	320	7095
	101	14		2180	4305	4603	-2125	298	6110
	102	14		2128	4203	4470	-2075	263	6062
	103	35		2154	4175	4420	-2021	245	3596
164	79	32		1492	2048	2320	-556	272	3775
	80	33		1508	2125	2470	-617	355	6874
	81	33		1519	2212	2576	-693	364	4780
	82	35		1534	2332	2670	-798	338	3706
	83	33		1603	2555	2877	-952	355	3051
	84	30		1643	2685	3098	-1042	413	7684
	85	33		1680	2790	3155	-1110	365	5388
	86	36		1737	2990	3360	-1253	370	3906
	87	34		1636	3055	3395	-1419	340	6349
	89	32		185 2	3452	3748	-1600	306	5332
	90	33		1901	3665	4017	-1764	352	5908
	91	31		1959	3775	4130	-1816	355	1707
	92	34		1978	3830	4195	-1852	365	2856
	93	31		1911	3915	4282	-2004	367	1153
	95	31		1903	3878	4248	-1975	370	5989
	96	32		1898	3982	4320	-2084	338	4149
	103	34		2177	4110	4315	-1933	205	3982

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APPENDIX C

VALUES OF THICKNESS FOR MEMBERS "A", "B", AND "C"

SE ST 1

Values of thickness for members "A", "B", and "C"

Appendix C consists of Table 2, which is a list of the wells used in the isopach maps of members "A", "B", and "C" (Plates X-XII). The wells are arranged by township (TWN), range (RNG), and section (SEC) and include elevation of the Kelly Bushing (KB), depth to the tops of members "C" (DC), "B" (DB), "A" (DA), and Swift Formation (DJSW), thicknesses of "C" (TC), "B" (TB), and "A" (TA), and the North Dakota Geological Survey well number (NDGS).

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TABLE 2

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<u>TWN</u> 146	RNG 96 97 98	SEC 10 8 36	<u>KB</u> 2658 2509 2262	DC 5713 5601 5368	<u>DB</u> 5743 5634 5399	<u>DA</u> 5824 5725 5454	<u>DJSW</u> 6117 5986 5666	TC 30 33 31	<u>TB</u> 81 91 55	<u>TA</u> 293 261 212	NDGS 7477 6335 6324	
	99 100	11 21	2628 2590	5791 5718	5836 5765	5931 5830	6169 6072	45 47	95 65	238 242	7611 6786	
147	92 93 95 96 97 98 99 100	10 8 23 22 18 36 27 34	2048 2212 2379 2625 2055 2513 2587 2525	4903 5072 5510 5703 5120 5586 5729 5274	4930 5100 5542 5735 5152 5618 5765 5320	4982 5219 5653 5823 5219 5676 5827 5364	5312 5483 5910 6066 5452 5940 6080 5568	27 28 32 32 32 32 36 46	52 119 111 88 67 58 62 44	330 264 257 243 233 264 253 204	7745 4957 7576 6623 8486 6345 7636 8341	
148	90 95 96 97 98 100	1 18 10 20 36 9	2064 2441 2095 2443 2206 2311	4773 5478 5072 5552 5327 5594	4792 5515 5104 5585 5361 5636	4820 5556 5166 5638 5405 5702	5160 5848 5428 5900 5675 5924	19 37 42 33 34 42	28 41 62 53 44 66	340 29 2 262 262 270 222	4375 2352 7614 3178 6597 6524	
149	90 91 92 95 96 97 98 99	5 22 33 24 16 24 2 11 10	1989 2102 2363 2149 2294 2499 2212 2232 2127	4800 4942 5187 5047 5210 5372 5253 5538 5495	4817 4959 5215 5072 5242 5402 5285 5570 5529	4857 5020 5347 5160 5278 5446 5350 5657 5590	5168 5322 5595 5422 5576 5750 5625 5828 5868	17 17 28 25 32 30 32 32 32 34	40 61 132 88 36 44 65 87 61	311 302 248 262 298 304 275 171 278	1194 793 7214 607 5936 6368 4085 3157 6493	
150	90 92 93 95 95 96 97 98 99	1 20 4 33 19 20 3 23 15	2190 2143 2198 2334 2342 2327 2244 2022 2141	4981 5000 5147 5163 5100 5152 5147 5293 5522	5000 5028 5175 5200 5130 5185 5180 5326 5558	5084 5097 5229 5264 5175 5228 5243 5400 5643	5360 5357 5522 5560 5484 5532 5571 5708 5960	19 28 37 30 33 33 33 36	84 69 54 45 43 63 74 85	276 260 293 296 309 304 318 308 317	7783 7457 4113 3731 6608 6849 7743 7704 7873	
151	90 94 95 96 97 99	13 10 13 34 11 27	2115 1956 2107 2432 2291 2278	4761 4744 4734 5081 5093 5629	4781 4775 4762 5113 5127 5662	4934 4833 4806 5160 5181 5701	5171 5152 5130 5457 5478 6055	20 31 28 32 34 33	153 58 44 47 54 46	237 319 324 297 297 347	4392 4594 3056 4095 7008 8092	
152	90	30	1881	4568	4581	4737	4942	13	156	205	4747	

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<u>TW</u>	$\frac{N}{2} \frac{RNG}{O3}$	SEC	<u>KB</u>	DC 5004	DB 5030	$\frac{DA}{5008}$	DJSW	$\frac{TC}{26}$	$\frac{TB}{68}$	TA	NDGS
	2 95 94	7	2190	4803	4833	4885	5190	20 30	52	305	3387
	95	7	2363	4865	4896	4961	5249	31	65	288	4013
	90	14 31	2451	4957	4989	5041	5348	32 33	53	307	6323 28110
		<u> </u>		50.0		2115		55	50	290	2045
15	3 92	25	2307	5222	5248	5330	5636	26	82	306	3317
	95	36	2379	4927	4045	5008	5345	29 30	05 51	335	4239
	96	10	2097	4475	4505	4570	4933	30	65	363	4494
	97	16	2052	5338	5372	5406	5810	34	34	404	7233
	33	1	2342	5102	2139	5700	0100	31	29	404	[233
15	4 90	22	2235	4823	4846	4896	5221	23	50	325	7414
	91	5 12	2305	5106	5129	5147	5484	23	18	337	6695
	95	18	2052	4297	4328	4368	4750	25 31	35 40	340	6184
	96	1	1955	4240	4270	4330	4702	30	60	372	1266
	99	36	2256	5630	5664	5712	6125	34	48	413	7074
15	5 90	21	2317	4835	4854	4880	5212	19	26	332	7220
	91	27	2315	5032	5052	5080	5450	20	28	370	6834
	93	15	2300	2313	5341 4764	4814	5732	28	63 50	328	2273
	95	18	2305	4524	4556	4587	4972	32	31	385	6362
	97	14	2253	4938	4970	5037	5362	32	67	325	7089
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	94	20	2331	4628	5028 4655	5071 11600	5351	28 27	43 1) 11	280	7741
	96	36	2335	4585	4618	4681	5015	33	63	334	5427
	97	27	2272	5045	5082	5145	5500	37	73	355	547
	99	łU	2201	5475	5511	5500	5904	30	49	344	3406
15'	7 91	4	2372	4946	4965	5015	5340	19	50	325	4682
	93	33	2271	4580	4608	4662	5022	28	54	360	6111
	90	2	2296	5111	5148	5187	5506	33 37	20 29	303	7789
	98	20	2213	5323	5358	5408	5717	35	50	309	3449
15	8 91	19	2379	5017	5040	5105	5370	23	65	265	5786
-	94	7	2368	4888	4917	4988	5283	29	71	295	5350
	95 06	21	2513	4909 1882	4941	5014	5329	32 30	73	315	5935 7565
	97	17	2276	5126	5160	5218	5533	34	58	315	8239

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<u>TWN</u> 158	<u>RNG</u> 99	SEC 3	<u>KB</u> 2150	<u>DC</u> 4942	<u>DB</u> 4974	5012	DJSW 5328	<u>TC</u> 32	<u>TB</u> 38	<u>TA</u> 316	<u>NDGS</u> 3252	
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160	90 91 92 93 94 95 96 97 98 99	30 5 21 23 5 35 19 10 24	2336 2414 2358 2371 2379 2372 2290 2307 2243 2236	4631 4658 4750 4805 4816 4662 4745 4865 4865 4722 4775	4655 4680 4772 4831 4841 4690 4770 4894 4753 4810	4725 4741 4820 4895 4898 4772 4835 4948 4809 4865	4978 4999 5132 5163 5198 5078 5083 5228 5040 5070	24 22 26 25 28 25 29 31 35	70 57 48 64 57 82 52 55 55	253 258 312 268 300 306 248 280 231 205	6028 5721 7917 2684 4101 5192 5009 5547 5248 7116	
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162	90 91 92 93 95 96 98	32 34 15 36 20 16 5	1964 1975 1948 2096 2136 2141 2221	3996 4103 4039 4364 4303 4388 4388 4482	4010 4121 4051 4390 4329 4410 4503	4044 4168 4105 4454 4402 4476 4562	4303 4410 4413 4726 4668 4728 4853	14 18 12 26 26 22 21	34 47 54 64 73 66 59	259 242 308 282 266 252 291	7664 7859 6081 6460 4074 6798 1900	

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APPENDIX D

CORE DESCRIPTION

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APPENDIX D

Core description

Description of core from the Matthew Iverson 1, NDGS 165, Sec. 1, Twn. 155 N., Rng. 96 W. was arranged as follows: rock lithology; rock color; rock constituents; sedimentary structures; miscellaneous. Terminology used in the descriptions are from Pettijohn, Potter, and Siever (1973) and Reineck and Singh (1980). Depths in feet (m) were taken from core-box labels as filed with the Core and Sample Library of the North Dakota Geological Survey, Grand Forks, North Dakota.

APPENDIX D

CORE DESCRIPTION

NDGS 165

Core Depth

Core Description

INYAN KARA FORMATION

Member "C" (in part):

4590-4592 Interbedded shale and siltstone; shale is black; thinly (1399-1399.6) Interbedded shale and siltstone; shale is black; thinly laminated; siltstone is light gray; thinly laminated to lensoidal.

Member "B":

4592-4632 Interbedded siltstone and shale; shale is black; thinly
(1399.6-1411.8) laminated; with abundant plant fragments near the base
that decrease in abundance up section; siltstone is
light gray; quartzose; in thin to thick laminae that
decrease in abundance up section; structures include
load structures (throughout core), erosional truncation
(Fig. 22), micro-dikes (Fig. 20), convoluted bedding
(Fig. 21), rip-up clasts (Fig. 19), and both horizontal
and vertical burrows (Figs. 19 and 20); bedding is
predominantly wavy and parallel but may be either
even and parallel or wavy and nonparallel.

Member "A":

4632-4647 Siltstone and shale; shale is light to dark gray; thinly (1411.8-1416.4) laminated; with ironstone nodules (Fig. 15), iron-oxide staining, and abundant plant fragments; siltstone is light gray; quartzose; occurs as burrow infilling or thin laminae.

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4648-4929 Missing section (not cored) (1416.7-1502.3)

Core Description

4930-4939 Coal, sandstone, siltstone and shale; core broken into (1502.7-1505.4) fragments; sandstone is quartzose, coarse-grained, with a dusty red-brown coating (possibly iron-oxide), with infilling of a few pores with a soft, white mineral (possibly gypsum), well indurated, often containing fragments of coal (Fig. 16); siltstone is quartzose and occurs as laminae interbedded with gray shale.

4940-4943
(1505.7-1506.6)
Sandstone; light gray; quartzose, fine-grained,
friable; no bedding evident, but mottling suggests some bioturbation.

4943-4964 Sandstone; medium gray; medium-grained,
(1506.9-1514.2) dominantly quartz plus black shale fragments with laminae to thin beds of light gray shale, friable; ripple cross-lamination and climbing ripples (Figs. 13 and 14).

4964-4968Sandstone; medium gray; medium-grained, dominantly(1513-1514.2)quartz plus black shale fragments with thin beds
of quartzose sandstone, friable; cross-bedded.

4968-4971 Sandstone; medium gray; coarse- to medium-grained, (1514.2-1515.1) dominantly quartz and black shale fragments, "white specks" (possibly altered feldspar), and stringers of dark, possibly carbonaceous, material, friable; fining-upwards sequence with cross-bedding.

4971Shale-pebble conglomerate; black; pebbles are rounded,(1515.1)calcareous, with a matrix of coarse-grained, quartzose
sandstone.

4971-4972 Sandstone; dark gray; coarse-grained, dominantly quartz (1515.1-1515.4) and black shale fragments with "white specks" (possibly altered feldspar), friable; cross-bedded.

Swift Formation (in part):

Core Depth

4972-4980 Shale; black; calcareous with crystals of pyrite; (1515.4-1518) thinly laminated; upper contact is abrupt.

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REFERENCES CITED

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PLATE I

SCALE IN MILES

LOCATION OF CROSS-SECTIONS A-A' TO I-I'

T1983 W266 platel

WITH WELL REFERENCE NUMBERS

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PLATE I

BRAD L. WARTMAN



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NDGS 7313

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SERIES	GROUP	FORMATION	MEM- BER	B		DLL		
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LOG TYPES		
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DLL Dual Laterolog	NDGS 6687	NDGS
LL-3 Laterolog-3		Confor
LL-7 Laterolog-7	~~~~~	Uncon

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LEGEND

14 k-up curve Footage from kelly bushing 1 referènce no S well.no. (Miles) Distance between wells

PLATE II CORRELATIONS OF THE MOWRY-INYAN KARA INTERVAL IN WESTERN NORTH DAKOTA

(For cross-section location see Plate I) Stratigraphic Datum: Top of Mowry Fm.

ormable formation contact(s)

onformable(?) Formation Contact

50-E

(MILES) (KILOMETERS)



19 NDGS 8020

SERIES GROUP FORMATION BER C COLORADO BELLE FOURCHE (in part) UPPER CRET. 7.8 12.5 MOWRY S NEWCASTLE 0 Ш C A SKULL CREEK A -0 Ш E X " C " 0 A "В" E ш IN Y A N K A R A 3 0 ____ _____ -3-UPPER JURASSIC SWIFT (in part)



R Gamma Ray		Back-up curve
Spontaneous Potential	28	Well reference no
-3 Laterolog-3	NDGS 6687	NDGS Well no.
L Dual Laterolog		Conformable form
		Unconformable(?)
	(miles) (kilometers)	Distance between we

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PLATE III CORRELATIONS OF THE MOWRY-INYAN KARA INTERVAL IN WESTERN NORTH DAKOTA

nation contact(s)

Footage from kelly bushing

formation contact

100-30 (METER

(MILES) (KILOMETERS)

(For cross-section location see Plate I) Stratigraphic Datum: Top of Mowry Fm.

> PLATE III Brad L. Wartman 1983

T1983 W266 Plate3



48 NDGS 8888

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SERIES GROUP FORMATION MEM UPPER CRET. COLORADO BELLE FOURCHE (in part) 6.8 MOWRY S 0 0 NEWCASTLE ш A 0 SKULL CREEK F A -0 Ш £ X 0 "C" A £ M Ш X INYAN KARA MA 0 NMM ____ ------SWIFT UPPER (in part)

> LEGEND LOG TYPES Back-up curve GR Gamma Ray 56 Well reference no. SP Spontaneous Potential DLL Dual Laterolog NDGS 6687 NDGS well no. Conformable formation contact(s) (miles) Distance between wells LL-3 Laterolog-3 LL-7 Laterolog-7 ------ Unconformable(?) formation contact

Z

(For cross-section location see Plate I) Stratigraphic Datum: Top of Mowry Fm.

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Footage from kelly bushing

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CORRELATIONS OF THE MOWRY-INYAN KARA INTERVAL IN WESTERN NORTH DAKOTA

100-30 (METEL

(MILES) (KILOMETERS)

> PLATE IV Brad L. Wartman 1983

T1983 W266 Plate 4




plate 5 Wall

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BRAD L. WARTMAN



PLATE VIII STRUCTURE CONTOUR MAP OF THE TOP OF THE INYAN KARA FORMATION

LEGEND

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- OUTLINE OF STUDY AREA
- O CONTROL POINT

DEPRESSION

LOCAL COMPLEX STRUCTURE (Red Wing Creek) C.I.=100 feet

Datum is mean sea level



PLATE VIII BRAD L. WARTMAN 1983

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PLATE IX **ISOPACH MAP OF THE** LEGEND INYAN KARA FORMATION C.I.=50 feet FORMATION THICKNESS GREATER THAN 400 FEET OUTLINE OF STUDY AREA CONTROL POINT 20 Miles 30 Km 0

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PLATE IX BRAD L. WARTMAN 1983

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Brad L. Wartman

1983



R. 100 W.

PLATE XI **ISOPACH MAP OF MEMBER "B"**

(For location of study area see Plate I) T1983 W266 plate 11 C.I.=10 feet

EGEND	
-60-	CONTOUR
	MEMBER T

0

BER THICKNESS GREATER THAN 60 FEET

CONTROL POINT

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PLATE XI BRAD L. WARTMAN 1983



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PLATE XII BRAD L. WARTMAN 1983