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Ground-water Movement in the Upper Glacial Aquifer in the Manorville Area, Town of Brookhaven, Long Island, New York, in November 1983

David A. Eckhardt USGS

Eliezer J. Wexler USGS

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GROUND-WATER MOVEMENT IN THE UPPER GLACIAL AQUIFER IN THE MANORVILLE AREA,
TOWN OF BROOKHAVEN, LONG ISLAND, NEW YORK, IN NOVEMBER 1983
by David A. V. Eckhardt and Eliezer J. Wexler

U.S. GEOLOGICAL SURVEY

Water-Resources Investigations Report 85-4035

Prepared in cooperation with TOWN OF BROOKHAVEN



Syosset, New York

UNITED STATES DEPARTMENT OF THE INTERIOR DONALD PAUL HODEL, Secretary

GEOLOGICAL SURVEY

Dallas L. Peck, Director

For additional information write to:

U.S. Geological Survey 5 Aerial Way Syosset, New York 11791 Telephone: (516) 938-8830 Copies of this report may be purchased from:

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CONVERSION FACTORS AND ABBREVIATIONS

Multiply inch-pound units	<u>by</u>	To obtain metric units
	Length	
<pre>inch (in) foot (ft) mile (mi)</pre>	25.40 0.3048 1.609	millimeter (mm) meter (m) kilometer (km)
	Flow	
foot per day (ft/d)	0.3048	meter per day (m/d)
	Specific capacity	
<pre>gallon per minute per foot [(gal/min)/ft]</pre>	0.2070	liter per second per meter [(L/s)/m]
	Hydraulic conductivity	-
foot per day (ft/d)	0.3048	meter per day (m/d)

GROUND-WATER MOVEMENT IN THE UPPER GLACIAL AQUIFER IN THE MANORVILLE AREA, TOWN OF BROOKHAVEN, LONG ISLAND, NEW YORK, IN NOVEMBER 1983

By David A. V. Eckhardt and Eliezer J. Wexler

Abstract

Water levels in 52 wells near the Manorville scavenger-waste-disposal facility in the Town of Brookhaven were measured in November 1983 to determine hydraulic gradients and the velocity of ground-water flow in the upper glacial aquifer. Ground water moves south-southeastward (S22°E) from the ground-water divide, about 6 miles north of the facility, to discharge points near East Moriches and beneath Moriches Bay. The hydraulic gradient beneath the disposal facility is 6.5 feet per mile (0.0012 foot per foot), and the rate of horizontal flow is 0.9 to 1.7 feet per day. Vertical movement of ground water since closure of the disposal facility in 1982 is probably negligible because the vertical gradients are small and the upper glacial aquifer is anisotropic. During operation of the facility, however, ground-water mounding may have developed beneath the unlined settling basins, which could have induced downward movement of water in the upper glacial aquifer.

INTRODUCTION

Ground water is the sole source of water supply for nearly 3.1 million residents of Nassau and Suffolk Counties on Long Island. The Town of Brookhaven, in central Suffolk County (fig. 1), is expected to have continued population growth; thus proper management of the ground-water resource is essential. The aquifer system that supplies water to the Town of Brookhaven consists primarily of Pleistocene sand and gravel, which is called the upper glacial aquifer, and a deeper sequence of Cretaceous sand, silt, and clay deposits that rests on bedrock.

In recent years, leachate from solid- and liquid-waste-disposal sites in the Town of Brookhaven has entered the upper glacial aquifer, but neither the extent of contamination nor the potential for further migration are known. In 1981, the U.S. Geological Survey, in cooperation with the Town of Brookhaven, began a hydrogeologic study to evaluate the movement of ground water at a scavenger-waste-disposal facility near Manorville, in the east-central part of the Town of Brookhaven (fig. 1). The facility received solid waste, liquid sewage, and sludge from 1964 through 1982. As a preliminary step in the study, the Geological Survey installed 19 wells in 1983 to evaluate the rates and direction of ground-water movement within the 35-acre disposal site and surrounding area.

Purpose and Scope

This report describes the observation-well network that was established in the upper glacial aquifer in the Manorville area, lists the water levels measured in November 1983, and describes the rates and direction of ground-water movement in the upper glacial aquifer in the area. It also presents two water-table maps--one of the Manorville area (36 mi 2 , pl. 1) and one of the waste-disposal-facility area (4 mi 2 , pl. 2).

Acknowledgments

The authors thank Elaine McKibbin, Director of Sanitation, Town of Brookhaven, and Joseph Lapienski, Manorville facility supervisor, for their cooperation and assistance during the study. The authors also thank Paul E. Moore, of McLean and Associates, Brookhaven, N.Y., for his help in coordinating the vertical control surveys.

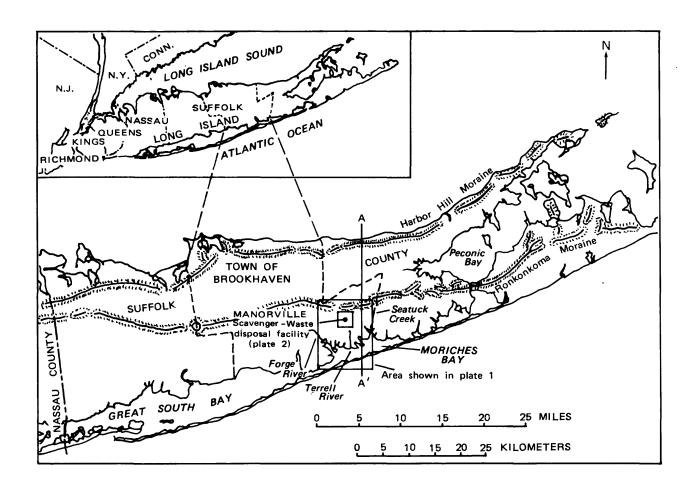


Figure 1.--Location of the Manorville area on Long Island. (Geologic section A-A' is shown in fig. 2.)

METHODS OF INVESTIGATION

A network of observation wells screened in the upper glacial aquifer was established in the Manorville area (pl. 1) and at the disposal facility (pl. 2) to define the ground-water flow patterns. First, an inventory of all wells in the 36-mi² Manorville area was made, and 32 wells were selected for the network. An additional 19 wells were then installed in the glacial outwash to supplement the network at the disposal site and surrounding areas. A reference-point altitude for each well was established by vertical control surveys from established benchmarks. The locations, physical characteristics, ownership, and other pertinent information about the wells are given in table 1. The wells range from 2 to 14 inches in diameter and are constructed of steel or polyvinyl chloride (PVC). A 14-inch observation well, owned by the Suffolk County Water Authority (S33920), is screened in the Magothy aquifer at 629 ft below land surface (pl. 2) and provides data for assessing vertical movement of ground water between the two aquifers.

Two or three observation wells were installed adjacent to each other at seven locations. Each well was screened at a different depth in the upper glacial aquifer to provide information about vertical head gradients and associated vertical flow of ground water between the screened intervals at each location. The well clusters can also be used for water-quality monitoring to provide data on the vertical distribution of contaminants.

Water levels in the wells were measured over a 6-hour period on November 23, 1983, and two water-table maps were constructed from the measurements. One is a regional map (pl. 1) that depicts general ground-water flow patterns in the Manorville area; the other (pl. 2) depicts flow patterns at the waste-disposal facility.

DESCRIPTION OF THE MANORVILLE AREA

The Manorville area, in the southeast part of the Town of Brookhaven (fig. 1), is bordered by the Peconic River on the north, Forge River on the west, Seatuck Creek on the east, and the saltwater of Moriches Bay and the Atlantic Ocean on the south (pl. 1). The surrounding region, known locally as the Pine Barrens, contains sandy soil that supports a natural growth of pitch pine and scrub oak. Irrigated cropland in the area supports prosperous vegetable farms and nurseries.

Climate

The climate in the Manorville area is influenced more by the ocean than by the adjacent mainland, and diurnal temperature changes are moderated accordingly. Precipitation, the only source of freshwater in this area, averages about 46 inches per year (Miller and Frederick, 1969) and is distributed fairly evenly throughout the year. Precipitation readily penetrates the sandy soil and glacial outwash, and direct runoff to streams is negligible. Evapotranspiration is about 22 inches, or nearly half the annual precipitation. Rates of evapotranspiration vary during the year and

Table 1.--Data on observation wells in the Manorville area.
[Locations are shown on pl. 1]

Well Depth no. Latitude/Longitude Ownerl Ownerl Depth (ft) Top (ft) S31460 405004 724659 TOB 100 90 S31461 404956 724642 TOB 68 62 S31462 405000 724643 TOB 73 67 S33919 404908 724730 SCDHS 78 67 S33920 404908 724730 SCDHS 629 588 S36150 405117 724903 USGS 45 43 S36152 405014 724438 USGS 66 62	Bottom (ft) 100 68 73 78 629 45	Well diameter (in) 4 4
no. Latitude/Longitude Owner ¹ (ft) (ft) S31460 405004 724659 TOB 100 90 S31461 404956 724642 TOB 68 62 S31462 405000 724643 TOB 73 67 S33919 404908 724730 SCDHS 78 67 S33920 404908 724730 SCDHS 629 588 S36150 405117 724903 USGS 45 43	(ft) 100 68 73 78 629	(in) 4 4 4
S31460 405004 724659 TOB 100 90 S31461 404956 724642 TOB 68 62 S31462 405000 724643 TOB 73 67 S33919 404908 724730 SCDHS 78 67 S33920 404908 724730 SCDHS 629 588 S36150 405117 724903 USGS 45 43	100 68 73 78 629	4 4 4
S31461 404956 724642 TOB 68 62 S31462 405000 724643 TOB 73 67 S33919 404908 724730 SCDHS 78 67 S33920 404908 724730 SCDHS 629 588 S36150 405117 724903 USGS 45 43	68 73 78 629	4 4
S31462 405000 724643 TOB 73 67 S33919 404908 724730 SCDHS 78 67 S33920 404908 724730 SCDHS 629 588 S36150 405117 724903 USGS 45 43	68 73 78 629	4 4
S33919 404908 724730 SCDHS 78 67 S33920 404908 724730 SCDHS 629 588 S36150 405117 724903 USGS 45 43	78 629	
\$33920 404908 724730 SCDHS 629 588 \$36150 405117 724903 USGS 45 43	629	,
\$36150 405117 724903 USGS 45 43		4
	45	14
S36152 405014 724438 USGS 66 62		2
	66	2
S46546 405131 724557 USGS 123 120	123	2
S46913 404920 724846 SCDHS 20 10	15	6
S46914 404917 724845 SCDHS 22 6	11	6
S46966 404952 724705 SCDHS 86 72	82	6
S47755 405136 724645 SCDHS 58 45	55	4
S48946 405121 724906 SCDHS 45 31	41	6
S65604 404928 724835 USGS 56 51	56	2
S67198 404959 724645 SCDHS 97 90	95	4
\$73790 405007 724648 USGS 61 58	61	2
S73791 404957 724655 USGS 61 58	61	2
\$73792 404959 724651 USGS 61 58	61	2
\$73793 405000 724643 USGS 56 53	56	2
\$73794 404945 724652 USGS 73 70	73	2
S73795 404946 724644 USGS 47 44	47	2
\$73796 404947 724639 USGS 55 52	55	2
\$73797 404949 724635 USGS 57 54	57	2
\$73798 404953 724623 USGS 66 63	66	2
\$73799 404956 724642 USGS 101 98	101	2
\$73800 405025 724622 USGS 83 80	83	2
\$73801 404929 724622 USGS 50 47	50	2
\$73802 404837 724848 USGS 34 30	34	2
\$73803	45	2 2
	22	
\$73805 404850 724527 USGS 53 50 \$73806 404931 724511 USGS 66 60	53 63	2 2
	98	2
\$73807 405102 724506 USGS 100 95 \$73808 405053 724826 USGS 60 54	57	2
\$73809 404937 724642 private ² 50 40	50	6
•		
\$73810 404923 724638 private ² 50 40 \$73811 405014 724657 TOB 85 80	50 85	6 4
\$73812 405014 724657 TOB 65 60	65	4
\$73813 404956 724645 TOB 88 83	88	4
S73814 404956 724645 TOB 68 63	68	4
\$73815 404956 724645 TOB 53 48 \$73816 404955 724652 TOB 70 65	53 70	4 4
\$73817 404955 724652 TOB 90 85	70 9 0	4
S73818 404955 724652 TOB 110 105	110	4
\$73819 404745 724812 CMFD 42		6
\$73820 404756 724653 CMFD 46		6
S73821 404755 724544 EMFD 45		6
\$73822 404828 724441 EMFD 45		6
\$73823 405009 724748 private ³ 70		3
\$74293 405017 724950 USGS 71 67	71	2
\$74294 405213 724811 USGS 36 32	36	2
\$74295 405045 724726 USGS 56 52	56	2

¹ TOB, Town of Brookhaven SCDHS, Suffolk County Department of Health Services

USGS, U.S. Geological Survey CMFD, Center Moriches Fire Department EMFD, East Moriches Fire Department

 $^{^2}$ Unused irrigation well.

 $^{^3}$ Unused domestic well.

are highest during summer (Warren and others, 1968). Precipitation recharges the ground-water system only when evapotranspiration losses are small and after soil-moisture deficiencies have been satisfied. On Long Island, the highest rates of recharge occur during late fall, winter, and early spring.

Hydrogeologic Setting

The hydrogeology of eastern Long Island has been described in detail by several authors, including Krulikas (1986), Koszalka (1984), Jensen and Soren (1974), McClymonds and Franke (1972), Warren and others (1968), and deLaguna (1963). A generalized geologic section of the aquifer system is shown in figure 2.

The Manorville disposal facility is excavated in highly permeable Pleistocene sand and gravel outwash that forms the upper glacial aquifer. The outwash is estimated to be about 160 ft thick at the facility and probably is underlain by the Gardiners Clay of Pleistocene age (fig. 2). Geologic information about clay units in the area is lacking, however.

The Monmouth Group, which consists of Cretaceous greensand and clay known as the Monmouth greensand, underlies the upper glacial aquifer in the southern part of the Manorville area. Its northern limit is uncertain but is assumed to be south of the disposal facility.

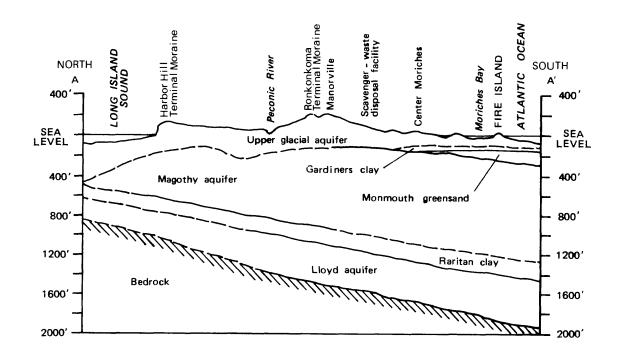


Figure 2.--Generalized hydrogeologic section showing relative positions of the major aquifers. (From Jensen and Soren, 1974.)

The Gardiners Clay and the Monmouth greensand, where present, unconformably overlie the Matawan Group and Magothy Formation, undifferentiated, of Cretaceous age, a thick body of continental deltaic deposits containing gradational sequences of sand, sandy clay, clay, and, in basal sections, some gravel. The Matawan-Magothy sequence may exceed 900 ft in thickness in the Manorville area and forms the Magothy aquifer.

The Raritan Formation of Cretaceous age, which underlies the Magothy Formation, consists of two members—an upper unnamed clay member that forms the Raritan confining unit, and the lower Lloyd Sand Member, which forms the Lloyd aquifer (fig. 2). The Raritan Formation rests on crystalline bedrock. The estimated thickness of the clay and sand member is 200 and 400 ft, respectively. The total thickness of unconsolidated material in the Manorville area, therefore, is more than 1,600 ft.

Hydraulic Characteristics of the Upper Glacial Aquifer

A compilation of pumping yields, drawdowns, and specific capacities (yield per foot of drawdown) reported by well drillers for 14 irrigation and public-supply wells in the Manorville area is given in table 2. Generally, these wells have a high specific capacity, with a median of 19 (gal/min)/ft. The median hydraulic conductivity in the upper glacial aquifer was computed from data in table 2 by a method described by McClymonds and Franke (1972, p. Ell) to be 430 ft/d. They indicate, however, that values calculated by this method may be higher than actual values because of vertical flow to partially screened wells. Thus, the value of 430 ft/d may represent an upper limit for horizontal hydraulic conductivity in this area.

An analysis of pumping tests in the upper glacial aquifer in southern Nassau County (Lindner and Reilly, 1983) indicated horizontal hydraulic conductivity of the glacial outwash to range from 140 to 380 ft/d. Mean hydraulic conductivity of the aquifer in the vicinity of the Horseblock Road sanitary landfill in the Town of Brookhaven, 8 mi west of the Manorville area (fig. 1), is estimated to be 220 ft/d (E. J. Wexler, U.S. Geological Survey, written commun., 1984). Because hydraulic conductivity controls the average rate of ground-water movement and associated transport of dissolved contaminants, an accurate estimate is essential for predicting contaminant movement. Although the values above give a probable range of values, additional aquifer testing would permit a more accurate determination of mean hydraulic conductivity.

Estimates of the ratio of horizontal to vertical hydraulic conductivity in the upper glacial aquifer range from 5:1 to 24:1 (Lindner and Reilly, 1983, p. 2). Porosity of the aquifer was assumed to be 30 percent (McClymonds and Franke, 1972).

Hydraulic Characteristics of the Magothy Aquifer

Few data on the hydraulic conductivity of the Magothy aquifer in the Manorville area are available because only one well in the area is screened in the formation. Regional estimates of horizontal hydraulic conductivity range between 13 and 54 ft/d (Warren and others, 1968) and from 40 to 54 ft/d

(McClymonds and Franke, 1972). Anisotropy in the Magothy aquifer (the ratio of horizontal to vertical hydraulic conductivity) was estimated by Getzen (1977) to range from 30:1 to 60:1. Aquifer porosity is assumed to be 25 percent (McClymonds and Franke, 1972).

Table 2.--Specific-capacity data from drillers' reports on wells screened in the upper glacial aquifer in the Manorville area.

[Well locations are shown in pl. 1]

Well	Depth	Screen length	Screen diameter	Pumping rate	Drawdown	Specific capacity [(gal/min)/	Specific capacity per ft of screen [(gal/min)/
no.	(ft)	(ft)	(in)	(gal/min)	(ft)	ft]	ft ²]
S11260	154	20	12	800	25	32	1.6
S20633	60	20	6	250	9	28	1.4
S21768	133	41	16	2,400	30	80	2.0
S23876	100	15	6	150	12	12.5	0.8
S34156	95	10	6	200	14	14	1.4
S35469	74	15	6	250	14	18	1.2
S46712	100	27	8	150	5	30	1.1
S46913	20	5	6	50	6	8.3	1.7
S46914	22	5	6	50	6	8.3	1.7
S46966	86	10	6	80	4	20	2.0
S47755	58	10	4	20	3	6.7	0.7
S48946	45	10	6	75	6	12.5	1.2
S56038	158	36	12	500	10	50	1.4
S56039	163	40	12	1,404	12	117	12.9

Well has 70-ft gravel pack, which effectively lengthens reported screen length and reduces specific capacity per foot of screen to about 1.7 (gal/min)/ft²

DESCRIPTION OF THE MANORVILLE WASTE-DISPOSAL FACILITY

The Manorville scavenger-waste-disposal facility is on Paper Mill Road in the southern part of Manorville (pl. 2). The facility covers 35 acres. It is at the head of the Terrell River valley, which contains flowing water intermittently, and lies on the glacial outwash plain immediately south of the Ronkonkoma terminal moraine (fig. 1). The site is bordered on the south and east by tree nurseries and on the north and west by woodland.

The Manorville scavenger-waste-disposal facility received municipal sludge and liquid septic wastes from domestic holding tanks for 9 years and may also have received liquid industrial wastes periodically. The wastes were deposited in a series of unlined settling basins that were excavated in the glacial outwash. The disposal facility had a design capacity of 50,000 gal/d, but incoming sewage and sludge volumes frequently exceeded that amount (Dvirka

and Bartilucci, 1981). The scavenger waste typically contained high amounts of suspended solids, both organic and inorganic, as well as high concentrations of total nitrogen, biochemical oxygen demand (BOD), and chemical oxygen demand (COD). Elevated concentrations of heavy metals and volatile organic compounds, particularly toluene, indicated some industrial waste in the scavenger waste (Dvirka and Bartilucci, 1981). The primary forms of waste treatment in the basins were chlorination and aeration of the fluids.

Liquid-waste disposal started in 1964 and continued until August 1982. Municipal solid wastes were also landfilled at the site during 1960-72. The facility was closed in November 1982 when it did not meet New York State environmental criteria for operation. After closure of the site, the basins were allowed to dry. Bottom material is currently (1984) being excavated and removed to a nearby sanitary landfill.

GROUND-WATER MOVEMENT

Observation wells provide a method for measuring water pressure at the point in the aquifer at which the well is screened. The depth to water below a known reference altitude indicates the total hydraulic head at that point, which is expressed in feet above sea level. Because water moves from areas of higher total head to areas of lower total head, directions of ground-water flow may be deduced from simultaneous head measurements over a given area. Estimates of ground-water velocities can be made from the gradients in hydraulic head and the hydraulic conductivity and effective porosity of the aquifer material.

Water-Level Measurements

The U.S. Geological Survey measured water levels in 52 observation wells on November 23, 1983. The water-table altitudes and the single Magothy aquifer potentiometric-surface measurement are given in table 3. The water-table contours within the Manorville area are shown on plate 1, and those at the waste-disposal facility on plate 2. Ground water flows perpendicular to water-table contours; thus, in the Manorville area, the direction of flow is south-southeastward (S22°E) from the ground-water divide near the Ronkonkoma moraine (fig. 1) toward East Moriches and Moriches Bay.

Rate of Flow

Ground-water flow rates were estimated from Darcy's law (Bear, 1979), which directly relates ground-water velocity to the hydraulic gradient between two points in an aquifer. The rate of ground-water flow in the aquifer was calculated by multiplying the hydraulic conductivity by the hydraulic gradient and then dividing by the effective porosity of the material (a measure of interconnected pore space through which water may move).

The hydraulic gradient that drives ground-water flow has two components-a horizontal gradient and a vertical gradient-which allows computation of horizontal and vertical ground-water flow rates. The horizontal (water-table) gradient in the upper glacial aquifer beneath the waste-disposal facility is

Table 3.--Measurement-point altitudes and ground-water levels in the Manorville area, November 1983.

[Well locations are shown on pls. 1 and 2]

well depth altitude water altitude (ft) no. (ft) (ft) above sea level 31461 68 68.23 43.49 24.74 31461 68 68.23 43.49 24.74 31462 73 70.45 45.21 25.24 33919 78 65.75 45.49 20.24 33920 629 66.10 46.13 19.97 36150 45 49.75 15.18 34.57 36152 66 64.68 43.68 21.00 46546 123 124.49 94.17 30.32 46914 22 32.98 22.60 22.60 46966 86 87.46 62.66 24.80 47755 58 59.74 27.40 32.34 48946 45 44.12 9.84 34.28 65604 56 64.18 40.17 24.01 73790 61 75.31			Measurement-point	Depth to	Water-level
no. (ft) (ft above sea level) (ft) above sea level 31460 100 89.14 63.25 25.89 31461 68 68.23 43.49 24.74 31462 73 70.45 45.21 25.24 33919 78 65.75 45.49 20.24 33920 629 66.10 46.13 19.77 36150 45 49.75 15.18 34.57 36152 66 64.68 43.68 21.00 46913 20 30.62 7.85 22.72 46914 22 32.98 22.60 22.60 46966 86 87.46 62.66 24.80 47755 58 59.74 27.40 32.34 48946 45 44.12 9.84 34.28 65004 56 64.18 40.17 24.01 73790 61 75.31 49.02 26.29 73791 61	Uo11	Well death			
31460 100 89,14 63,25 25,89 31461 68 68,23 43,49 24,74 31462 73 70,45 45,21 25,24 33919 78 65,75 45,49 20,24 33920 629 66,10 46,13 19,97 36150 45 49,75 15,18 34,57 36152 66 64,68 41,68 21,00 46964 62,66 123 124,49 94,17 30,32 46914 22 32,98 22,60 22,60 44966 86 87,46 62,66 24,80 47755 58 59,74 27,40 32,34 48946 45 44,12 9,84 34,28 65604 56 64,18 40,17 24,01 67198 97 79,35 54,05 25,30 73791 61 84,27 59,15 25,12 73792 61					•
31461 68 68.23 43.49 24.74 31462 73 70.45 45.21 25.24 33919 78 65.75 45.49 20.24 33920 629 66.10 46.13 19.97 36150 45 49.75 15.18 34.57 36152 26 64.68 43.68 21.00 46913 20 30.62 7.85 22.72 46914 22 32.98 22.60 22.60 44966 86 87.46 62.66 24.80 47755 58 59.74 27.40 32.14 48946 45 44.12 9.84 34.28 65604 56 64.18 40.17 24.01 67198 97 79.35 54.05 25.30 73791 61 75.31 49.02 26.29 73792 61 84.27 59.15 25.12 73793 56 68.77 43.59 25.31 73796 55 67.05 43.45 <td< td=""><td></td><td><u>\</u></td><td>(10 00010 000 10101)</td><td></td><td></td></td<>		<u>\</u>	(10 00010 000 10101)		
31461 68 68.23 43.49 24.74 31462 73 70.45 45.21 25.24 33919 78 65.75 45.49 20.24 33920 629 66.10 46.13 19.97 36150 45 49.75 15.18 34.57 36152 26 64.68 43.68 21.00 46913 20 30.62 7.85 22.72 46914 22 32.98 22.60 22.60 44966 86 87.46 62.66 24.80 47755 58 59.74 27.40 32.14 48946 45 44.12 9.84 34.28 65604 56 64.18 40.17 24.01 67198 97 79.35 54.05 25.30 73791 61 75.31 49.02 26.29 73792 61 84.27 59.15 25.12 73793 56 68.77 43.59 25.31 73796 55 67.05 43.45 <td< td=""><td>31460</td><td>100</td><td>89.14</td><td>63.25</td><td>25.89</td></td<>	31460	100	89.14	63.25	25.89
33919 78 65.75 45.49 20.24 33920 629 66.10 46.13 19.97 36150 45 49.75 15.18 34.57 36152 26 64.68 43.68 21.00 46913 20 30.62 7.85 22.72 46914 22 32.98 22.60 22.60 46966 86 87.46 62.66 24.80 47755 58 59.74 27.40 32.34 48946 45 44.12 9.84 34.28 65604 56 64.18 40.17 24.01 67198 97 79.35 54.05 25.30 73791 61 75.31 49.02 26.29 73792 61 80.30 54.99 25.31 73793 56 68.77 43.59 25.12 73794 73 85.56 61.83 23.73 73795 47 56.75	31461	68		43.49	24.74
33920 629 66.10 46.13 19.97 36150 45 49.75 15.18 34.57 36152 66 64.68 43.68 21.00 46546 123 124.49 94.17 30.32 46913 20 30.62 7.85 22.72 46914 22 32.98 22.60 22.60 46966 86 87.46 62.66 24.80 47755 58 59.74 27.40 32.34 48946 45 64.18 40.17 24.01 65604 56 64.18 40.17 24.01 67198 97 79.35 54.05 25.30 73790 61 75.31 49.02 26.29 33792 61 80.30 54.99 25.31 73792 61 80.30 54.99 25.31 73794 73 85.56 61.83 23.73 73795 57 56.75	31462	73	70.45	45.21	25.24
36150 45 49.75 15.18 34.57 36152 66 64.68 43.68 21.00 46546 123 124.49 94.17 30.32 46913 20 30.62 7.85 22.72 46914 22 32.98 22.60 22.60 46966 86 87.46 62.66 24.80 47755 58 59.74 27.40 32.34 48946 45 44.12 9.84 34.28 65604 56 64.18 40.17 24.01 67198 97 79.35 54.05 25.30 73790 61 75.31 49.02 26.29 73791 61 84.27 59.15 25.12 73792 61 80.30 54.99 25.18 73793 56 68.77 43.59 25.18 73794 73 85.56 61.83 23.73 73795 47 56.75 33.10 23.65 73796 55 67.05 43.45 23.60 73799 101 69.02 44.28 24.74 73800 83 101.11 74.26 26.85 73801 50 65.24 44.29 20.95 73803 48 45.91 32.48 13.43 73805 53 46.04 33.25 12.79 73806 66 50.86 35.15 15.71 73807 100 108.75 82.15 26.60 73810 50 37.84 17.22 73807 100 108.75 82.15 26.60 73810 50 37.84 17.22 73809 50 38.30 16.01 22.29 73810 50 37.84 17.22 73811 85 83.47 56.10 27.37 73812 65 83.47 56.10 27.37 73813 88 64.11 39.32 24.81 73815 53 64.17 39.35 24.82 73816 70 87.06 62.21 24.85 73817 90 87.06 62.21 24.85 73818 110 87.02 62.17 24.85 73819 20 13.12 10.03 3.09 73822 20 12.54 66.66 5.88 73822 20 17.18 12.14 5.04 73823 71 88.37 53.14 30.23	33919	78	65.75	45.49	20.24
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6.5 ft/mi or 0.0012 ft/ft. With this gradient, an average hydraulic conductivity of 220 to 430 ft/d, and a porosity of 0.30, the rate of horizontal flow ranges from about 0.9 to 1.7 ft/d. At these flow rates, ground water beneath the facility will move from 6,600 to 12,000 feet south-southeastward in the upper glacial aquifer in 20 years. The flow patterns and gradients at Manorville agree with those reported in a study of the nearby Pine Barrens area of Suffolk County (Krulikas, 1986). The flow patterns and gradients reported for the Manorville area represent measurements on only one date, however, and do not account for possible seasonal variations.

Vertical gradients were measured at the waste-disposal facility to determine rates of vertical movement of water to deeper zones within the upper glacial aquifer. Aquifer heads at five well clusters near the facility differed slightly and indicate downward movement of water, but the differences were extremely small--less than 0.01 ft--which is below the limits of accuracy for vertical-control surveys of measurement-point altitudes and water-level measurements. Because these vertical gradients were small, and also because the vertical hydraulic conductivity of the upper glacial aquifer may only be from 1/5 to 1/24 of the horizontal value, the vertical movement of ground water beneath the disposal facility since its closure in 1982 is probably negligible.

During operation of the disposal facility, mounding of the water table may have occurred as wastewater seeped out of the unlined basins. Although the regional ground-water movement in this area is predominantly horizontal, increased vertical gradients during mounding could have caused downward movement of wastewater within the aquifer. The water-table mounds would have dissipated shortly after waste-disposal operations ceased in 1982, however.

Measurable differences in vertical heads were observed between well \$33920, which is screened at an intermediate depth in the Magothy aquifer, and an adjacent well, \$33919, screened in the upper glacial aquifer (well locations are shown in pl. 2). The vertical gradient measured at this well cluster in November 1983 was equal to 0.0005 ft/ft and indicates that ground water moves downward from the upper glacial aquifer to the Magothy aquifer at this site. The rate of downward movement can be calculated from estimates of effective porosity, hydraulic conductivity, and the ratio of anisotropy for the Magothy aquifer. The computed rates of vertical movement are extremely low, between 0.16 and 1.3 ft/yr. If the Gardiners Clay overlies the Magothy aquifer beneath the disposal facility, rates of vertical movement would be considerably lower. Additional hydrogeologic data would be required to more fully appraise the potential for downward movement of water from the upper glacial aquifer to the Magothy aquifer.

SUMMARY AND CONCLUSIONS

Water-level measurements were made at 52 observation wells in the 36-mi² area near the Manorville scavenger-waste disposal facility in November 1983 to determine the direction and velocity of ground-water flow. Ground water moves south-southeastward (S22°E) from the ground-water divide north of the facility toward East Moriches and Moriches Bay. Horizontal flow at the

facility is induced by a 6.5-ft/mi (0.0012) hydraulic gradient on the water table. From an average hydraulic conductivity ranging from 220 to 430 ft/d and an effective porosity of 0.30, the average ground-water flow rate in the upper glacial aquifer beneath the disposal facility may range from 0.9 to 1.7 ft/d. Within this range of flow rates, ground water beneath the facility will move 6,600 ft to 12,000 ft south-southeastward in the upper glacial aquifer in 20 years. The principal direction of ground-water movement in the upper glacial aquifer is horizontally toward discharge points near East Moriches and beneath Moriches Bay.

Vertical gradients at the Manorville disposal facility are small compared to the average horizontal (water-table) gradient of 0.0012 ft/ft. Vertical movement of ground water beneath the disposal facility since its closure in 1982 is probably negligible because of the small vertical gradients and the low vertical hydraulic conductivity of the upper glacial aquifer (1/5 to 1/24 of the horizontal value). Mounding of the water table below the unlined settling basins during operation of the facility, however, may have caused downward movement of wastewater.

The vertical gradient measured at a well cluster 1.1 mi downgradient of the disposal facility was used to compute the rate of downward flow of ground water from the upper glacial aquifer to the Magothy aquifer. The estimated rates of downward flow are low, ranging from 0.16 to 1.3 ft/yr.

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