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Modelling Emergent Patterns of Dynamic Desert Ecosystems

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Modelling Emergent Patterns of Dynamic Desert Ecosystems

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ABSTRACT

1
 2 In many desert ecosystems vegetation is both patchy and dynamic: vegetated areas are
 3 interspersed with patches of bare ground, and both the positioning and the species composition
 4 of the vegetated areas exhibit change through time. These characteristics lead to the emergence
 5 of multi-scale patterns in vegetation that arise from complex relationships between plants, soils
 6 and transport processes. Previous attempts to probe the causes of spatial complexity and predict
 7 responses of desert ecosystems tend to be limited in their focus: models of dynamics have been
 8 developed with no consideration of the inherent patchiness in the vegetation, or else models have
 9 been developed to generate patterns with no consideration of the dynamics. Here we develop a
 10 general modelling framework for the analysis of ecosystem change in deserts that is rooted in the
 11 concept of connectivity and is derived from a detailed process-based understanding. We
 12 explicitly consider spatial interactions among multiple vegetation types and multiple resources,
 13 and our model is formulated to predict responses to a variety of endogenous and exogenous
 14 disturbances. The model is implemented in the deserts of the American Southwest both to test
 15 hypotheses of the causes of the invasion of woody shrubs, and to test its ability to reproduce
 16 observed spatial differences in response to drought in the twentieth century. The model's
 17 performance leads us to argue that vertical and lateral connectivity are key emergent properties
 18 of the ecosystem, which both control its behaviour and provide indicators of its state. If this
 19 argument is shown to be compatible with field observations, the model presented here will
 20 provide a more certain approach towards preventing further degradation of semi-arid grasslands.

21
 22 **Key words:** semi-arid landscape ecology; patchiness; vegetation patterns; positive feedback;
 23 spatially explicit model; self-organization

24

25 **1. Introduction**

26 Desert ecosystems are commonly dynamic and patchy on a range of spatial and temporal scales
 27 (Ward, 2008; Wainwright, 2009). Their dynamism is particularly evident in long-term data
 28 showing changes in the composition and structure of plant communities (Buffington and Herbel,
 29 1965; Gibbens *et al.*, 2005; Turner *et al.*, 2003; Osborne and Beerling, 2006; Schlesinger *et al.*,
 30 1990; Archer *et al.*, 1995; Allred, 1996). Their patchiness, in which vegetated areas are
 31 interspersed with areas of bare ground, varies with plant growth form. In grasslands, bare and
 32 grassy patches alternate over a few decimetres and, on sloping ground, are often associated with
 33 a stepped microtopography (Dunkerley and Brown, 1999; Tongway and Ludwig, 2001; Nash *et*
 34 *al.*, 2004; Parsons *et al.*, 1997). In shrublands, the spatial scale extends to a few metres and the
 35 microtopography may comprise swales (bare patches) and vegetation atop mounds (Barbour,
 36 1969; McPherson *et al.*, 1988; Parsons *et al.*, 1996; Rango *et al.*, 2000; Okin and Gillette, 2001).
 37 This patchiness in vegetation can lead to the formation of striking, regular patterns such as bands
 38 of vegetation alternating with stripes of bare patches (Gillett, 1941), exemplified by ‘tiger bush’
 39 in Africa (MacFadyen, 1950; Clos-Arceud, 1956), mulga groves in Australia (Slatyer, 1961)
 40 and mogote in Mexico (Cornet *et al.*, 1988). Other geometric and irregular patterns have also
 41 been noted such as can be described as spots (Bromley *et al.*, 1997) and labyrinths (Aguiar and
 42 Sala, 1999) in the dry zones of the world.

43

44 The formation of vegetation patches has typically been explained in two ways. On the one hand,
 45 the empirical-conceptual model of Islands of Fertility has been used since its definition by
 46 Charley and West (1975), and especially since its development by Schlesinger *et al.* (1990), to
 47 explain patches at the scale of individual plants. On the other hand, advection-diffusion models

48 – usually, but not always (e.g. Lefever *et al.*, 2009) of the Turing-instability type (e.g.
 49 Klausmeier, 1999; Couteron and Lejeune, 2001; HilleRisLambers *et al.*, 2001), have been
 50 employed to explain patterns at landscape scales.

51
 52 Both of the existing explanations of patchiness have shortcomings. Both have conceptual
 53 limitations, and they produce results that are mutually incompatible and difficult to evaluate
 54 independently. The Islands of Fertility approach attempts to consider the system dynamics, but it
 55 is poor at addressing the emergence of spatial patterns. Conversely, the advection-diffusion
 56 approach is able to simulate emergence of spatial patterns, but often at the expense of an
 57 appropriate characterization of the dynamics. The aim of this paper is to advance our
 58 understanding of the dynamics of desert ecosystems and the patchiness and patterns that result
 59 from these dynamics. Using a conceptual approach with an explicit process basis, a new model
 60 is developed that links the dynamics of desert ecosystems with vegetation patchiness that is both
 61 quantitative and testable against existing data.

62
 63
 64 **2. Existing approaches – a critique**

65 **2.1 The Islands of Fertility Model**

66 The Islands of Fertility model (Charley and West, 1975; Schlesinger *et al.*, 1990) posits that
 67 changes in the spatial redistributions of soil resources are caused by the net transport of resources
 68 from interspaces to under-canopy areas. The heterogeneous resource distribution in turn affects
 69 plant demographic processes to reinforce vegetation patchiness. For example, during rainstorms,
 70 patches of vegetation serve as obstructions that slow, trap and accumulate runoff, sediments and

71 nutrients from interpatch areas (Ludwig *et al.*, 2005). This accumulation leads to the increase of
 72 patch biomass, which will further accumulate resources (Aguiar and Sala, 1999). Establishment
 73 is reported to be particularly successful around the edges of the patch where there is less
 74 competition for sunlight (Mauchamp *et al.*, 1993), and where flows of nutrients and water
 75 become trapped, such as on the upslope edge of a patch (Montana, 1992). Where individual
 76 plants die, wind- and water-induced degradation of the patch are increased, which reduces seed
 77 establishment. Propagules are then moved to other locations where they may establish new
 78 communities (Goldberg and Turner, 1986).

79
 80 Over time, changes in concentrations of resources may lead to new vegetation species attaining a
 81 competitive advantage within these patchy ecosystems (Osborne and Beerling, 2006). For
 82 example, the widely observed encroachment of shrubs into former grasslands (Allred 1996,
 83 Schlesinger *et al.*, 1990, Archer *et al.*, 1995) is regarded as a process that, due to different spatial
 84 distributions of grasses and shrubs, results in self-reinforcing changes to the spatial
 85 redistributions of soil resources (Schlesinger *et al.*, 1996). However, Islands of Fertility cannot
 86 explain all the different scales at which patterns appear in desert vegetation (Müller *et al.*, 2008).
 87 Islands have also been demonstrated to be “leaky” (Wainwright *et al.*, 2002; Abrahams *et al.*,
 88 2002) and thus linked to ecogeomorphic processes occurring beyond the scale of individual
 89 islands. Furthermore, the Islands of Fertility model does not tell us how changes are initiated,
 90 simply why they persist. Although the model explains why invading shrubs have a competitive
 91 advantage, it does not explain how they were able to invade in the first place, nor why certain
 92 types of patterns occur (e.g., stripes) occur under some circumstances. A broader issue with the
 93 Islands of Fertility model is that the term itself is tautologous, and hence unsuitable for predictive

94 purposes. If concentrations of resource are present around a shrub, it is an island. If
 95 concentrations are absent around a shrub, it is not an island. Because of the qualitative and
 96 descriptive nature of the approach (Schlesinger *et al.*, 1990; Ludwig *et al.*, 2005), there is
 97 nothing independent of the resource accumulations that would allow this idea to be tested.
 98 Although Schlesinger *et al.* (1996) used semivariograms to support the idea of spatial patterns,
 99 these patterns are still not independent of the Islands that they are meant to demonstrate.

100

101 **2.2 Numerical Models**

102 The use of numerical modelling to shed light on dryland vegetation has, so far, been limited in
 103 terms of its narrow focus: models of dynamics have either been developed with no consideration
 104 of the inherent patchiness or patterns in the vegetation (Thornes and Brandt, 1993 [and
 105 discussion in Wainwright and Parsons, 2010]; Peters, 2002a; Koppel and Rietkerk, 2004;
 106 Istanbuluoglu and Bras, 2006), or they have been developed to generate patterns with no
 107 consideration of the dynamics (Klausmeier, 1999; Couteron and Lejeune, 2001; HilleRisLambers
 108 *et al.*, 2001; Rietkerk *et al.*, 2002; van de Koppel and Rietkerk, 2004; Barbier *et al.*, 2006), or
 109 they have been parameterized to create a specific ecosystem response (e.g. Dakos *et al.*, 2011).
 110 Furthermore, if, as seems widely believed, both dynamics, and patterns/patchiness and
 111 ecosystem responses are functions of resource (principally water) limitation, then there has been
 112 little integration into these models of the temporal and spatial variability of resource availability
 113 that are well documented for deserts (Noy-Meir, 1973; Comrie and Broyles, 2002; Wainwright *et*
 114 *al.*, 2000).

115

116 The most prevalent type of spatial model uses a Turing-like instability to generate regular
 117 patterns in desert vegetation. Patterns (Turing structures) originate solely through the coupling
 118 of reaction and diffusion processes, and the definition of a Turing structure specifically *excludes*
 119 any type of hydro-dynamic (i.e. fluid) motion (Turing, 1952). Despite contravening Turing's
 120 definition regarding applicability to hydrodynamic systems, this methodology was applied to
 121 vegetation patterns by Klausmeier (1999), whose model was based on the assumption that water
 122 cannot infiltrate on bare areas, so it flows downhill into a vegetation stripe where it does
 123 infiltrate and support plant growth. The flow of water was assumed to be exhausted before it
 124 reaches the downslope side of the stripe where the plants will consequently die off leading to a
 125 gradual uphill movement in the vegetation bands (Montana, 1992, but see above). Klausmeier's
 126 model involved the solution of two differential equations for water and plant biomass and
 127 assumed a uniform evaporation rate and water supply that is a linear function of increasing
 128 infiltration with increasing plant biomass. The model was reported to be insensitive to the exact
 129 form of functions of growth and infiltration as the resulting patterns are generated entirely by the
 130 Turing instability.

131
 132 The patterns in this type of model result from spontaneous symmetry-breaking phenomena
 133 associated with bifurcations of steady states, corresponding to stable stationary solutions to a set
 134 of reaction-diffusion equations (Nicolis and Prigogine, 1977; Meinhardt, 1982). In Klausmeier's
 135 work, terms for water supply, infiltration and growth represent the contributions of reactive
 136 processes while the diffusion terms, such as plant dispersal, bring in the spatial dependence. The
 137 reactive processes were set to give realistic values of the intrinsic relative periodicity of the
 138 resulting banded patterns. In doing so, however, highly implausible values for input parameters

139 had to be set; for example, water input of up to 750 mm a^{-1} and zero infiltration. Although some
 140 banded vegetation is found in areas with up to 750 mm a^{-1} , this is the exceptional (of the order of
 141 two to three times higher than the rate in areas where banded vegetation is typically observed),
 142 and observed infiltration rates are non-zero (see, for example Abrahams and Parsons, 1991;
 143 Casenave and Valentin, 1992), requiring an even more unrealistic rainfall input to match model
 144 output.

145
 146 Klausmeier's approach was extended by HilleRisLambers *et al.* (2001) and Rietkerk *et al.* (2002)
 147 so that the water input could be separated into a soil-water component. There are two major
 148 problems with the model of HilleRisLambers *et al.* (2001) and Rietkerk *et al.* (2002). First, the
 149 key conclusions drawn were that herbivory, plant dispersal, rainfall, drought intolerance and
 150 infiltration rate are not the primary factors that are likely to form patterns in vegetation.
 151 However, these factors are represented as the reactive processes in the Turing structures.
 152 Chandrasekhar (1961) and Klausmeier (1999) had already demonstrated that only the relative
 153 periodicity of patterns depends on factors controlling the reactive processes – the resulting
 154 patterns themselves are insensitive. Model output showing the formation of patterns in
 155 vegetation that are largely not controlled by the levels of water input, plant demographic
 156 characteristics and land-management practices is difficult to justify on the basis of our
 157 understanding of field processes. Experimental studies have identified these variables as being
 158 of significant importance (Coffin and Lauenroth, 1990; Parsons *et al.*, 1997, Parsons *et al.*,
 159 2006a; Parsons *et al.*, 2006b). Field observations have also noted that in some areas, there is an
 160 apparent relationship between rainfall and pattern type (e.g. Deblauwe *et al.*, 2008). However,
 161 this relationship is not universal, and very different patterns can be observed within an area of a

162 few square kilometres, which is too small to be explained by the existence of a precipitation
 163 gradient. The second problem is that the authors reported that without positive feedback
 164 between vegetation density and water infiltration, pattern formation was not found. However,
 165 this linkage is already defined as an essential condition required for Turing instabilities – that the
 166 kinetics should include a positive feedback process (Murray, 1989), so the result really states that
 167 without one of the essential conditions for Turing instabilities being included, Turing patterns do
 168 not form. Although this statement is mathematically true, it does nothing to further the
 169 understanding of pattern formation in vegetation.

170
 171 These more recent models are also applicable to and capable of generating patterns on flat
 172 surfaces, and it has subsequently been argued that this result invalidates the class of model that
 173 generates patterns only when some degree of pre-patterning is first applied (Couteron and
 174 Lejeune, 2001; Barbier *et al.*, 2006). However, all reported implementations of Turing-type
 175 models require some degree of pre-patterning. For example, Rietkerk *et al.* (2002) perturbed
 176 small amounts of plants or water in some areas of the simulation to generate patterns, and even in
 177 the original work of Couteron and Lejeune (2001), cells in their simulation were perturbed by a
 178 low level of noise. For a Turing-type model such pre-patterning effectively means that the
 179 model must produce a pattern. Moreover, according to Couteron and Lejeune (2001), there is no
 180 evidence in the literature of patterns appearing in arid or semi-arid environments devoid of a
 181 consistent source of anisotropy.

182
 183 The problems of the Turing-instability models highlight an important consideration that should
 184 be made for all numerical modelling techniques – if the underlying mathematical method is

185 designed to generate a pattern, a pattern will be generated. This outcome is reasonable when the
 186 mathematics describe a real process. For example, following previous field observations
 187 (Thornes, 1990) the model of Thornes and Brandt (1993) was set up to favour shrubs, and
 188 therefore showed a continued dominance of shrubs. By contrast, published Turing-type models
 189 of vegetation patterning do not rest on field observations, but rely on many parameters that
 190 would be difficult or impossible to measure in the field, such as ‘half saturation constant’, or are
 191 defined simply to produce the desired result (see comments above about rainfall rates).

192 Moreover, for Turing-type models, the underlying mechanism for symmetry-breaking requires
 193 coefficients to become negative under certain conditions – for example in the formulations
 194 described here, this would mean that there could be a negative water input for certain spacings of
 195 vegetation, which is meaningless in a physical sense. This inherent lack of realism in Turing
 196 patterns was reported by Rovinsky (1987), who noted that pattern formation could only occur
 197 where values of diffusion coefficients were in contradiction with physical arguments. Castets *et*
 198 *al.* (1990) and later Barbier *et al.* (2006) also observed that no unambiguous experimental
 199 observation of Turing patterns had been found. This ambiguity is drawn further into question by
 200 the contradictory conclusions drawn from Turing-instability models. For example,
 201 HilleRisLambers *et al.* (2001) argued that infiltration is not a primary causal factor in the
 202 formation of vegetation patterns, but van de Koppel and Rietkerk (2004) state that patterns do
 203 not form unless infiltration rate is linked to biomass. However, this linkage is not easy to define
 204 as the relationship between infiltration and biomass is not a straightforward one (Wainwright,
 205 2009).

206

207 Other, fundamental problems with applications of the Turing model have been revealed when
 208 modelling results are compared to field data. The implementation reported by Couteron and
 209 Lejeune (2001) more closely follows the classic description of Turing instabilities applied to
 210 chemical reactions where the diffusion term is alternatively expressed as an activator
 211 (propagation) and an inhibitory (competition) process. In this case, the essential condition for a
 212 Turing pattern is that the inhibitor must diffuse much more quickly than the activator (Castets *et*
 213 *al.*, 1990). In ecological terms, this condition would imply that the competition processes must
 214 occur over a larger scale than facilitative ones. While this difference in scale of operation may
 215 be appropriate for the consideration of a single species (e.g. a shrub canopy that facilitates
 216 growth by intercepting rainfall and channelling it to the roots operates over a smaller scale than
 217 that of competing root systems of adjacent plants: Abrahams *et al.*, 2003; Martinez-Meza and
 218 Whitford, 1996; Gibbens and Lenz, 2005; Brisson and Reynolds, 1994), it is not necessarily
 219 appropriate when species of different types are competing (e.g. as shrubs invade grassland, the
 220 competition effect of canopy interception by the shrub is at the same scale as any facilitation due
 221 to shading). Furthermore, the resulting relative periodicity observed by Klausmeier (1999) was
 222 noted to be much larger than the range of interactions between plants, and it has been argued that
 223 “somehow” local processes are amplified by the spatial instability (Castets *et al.*, 1990: 618).
 224 When the results of the model were compared to digitised images, Couteron and Lejeune (2001)
 225 found that the model yields much lower values for periodicity than were observed in field-site
 226 data. In their work, Couteron and Lejeune (2001) used realistic input data and found that the
 227 Turing model yields unrealistic intrinsic periods, whereas Klausmeier (1999) had used unrealistic
 228 input data to yield plausible intrinsic periods. A further problem arose when convincing field
 229 evidence was sought of patterns that are generated by Turing-type models, specifically the

230 hexagonal pattern that corresponds to bare spots appearing in continuous cover of annual grasses.
 231 No detectible hexagonal symmetry was found in the digitised images examined by Couteron and
 232 Lejeune (2001) nor in the subsequent work of Barbier *et al.* (2006). Many reasons were
 233 proposed to explain why this pattern could not be detected, except one – that the model had
 234 generated a pattern that does not exist in the real world because it was an inappropriate
 235 representation of the real world.

236
 237 An alternative to Turing-type models is provided by other instability-based approaches
 238 characterized by representations of short-range facilitation and long-range competition (e.g. Kéfi
 239 *et al.*, 2007; Lefever *et al.*, 2009). Whilst these alternative models overcome some of the
 240 limitations of the Turing-instability approach, they are nonetheless difficult to interpret in terms
 241 of physical processes (as discussed above).

243 **2.3 Beyond the limitations of existing approaches**

244 Both the Islands of Fertility model and those based upon instability approaches address patterns
 245 in desert vegetation that exist at a specific scale (plant and patch scale). However, there has been
 246 a recognition in recent years that there are patterns in desert vegetation that exist at larger scales,
 247 such as community and landscape scales (Wainwright *et al.*, 2002; Peters *et al.*, 2005; Turnbull
 248 *et al.*, 2008), which are addressed neither by the Islands of Fertility nor by instability-type
 249 models. The latter have been further hamstrung by the conceptual limitations of the models –
 250 applying models that are inappropriate and based on forcing process representations to fit a
 251 model structure rather than modelling the actual processes – and thus the practical limitations of

252 producing meaningful parameterizations. Both approaches are limited by the way in which
 253 inherent problems with definition prevent adequate testing.

254

255 Modelling studies are nevertheless useful for probing the causes and consequences of the
 256 observed dynamic patterns of desert vegetation. Not only may these studies provide a rigorous
 257 test of our understanding of ecosystem functioning, they can also provide detailed information
 258 that is not readily available from experimental work (for example, due to the timescales involved
 259 in measuring changes in plant communities [Parshall *et al.*, 2003] or due to ethical considerations
 260 such as large scale experimental modification of vegetation and landscapes). Recent conceptual
 261 advances into the understanding of multi-scale processes in drylands have focused on the rôle of
 262 process connectivity, drawing on broader concepts of landscape connectivity (Taylor *et al.*,
 263 1993; Turner *et al.*, 1993; Western *et al.*, 2001; Bracken and Croke, 2007; Müller *et al.*, 2007;
 264 Turnbull *et al.*, 2008; Okin *et al.*, 2009). Turnbull *et al.* (2008) suggested that spatial patterns
 265 emerge as the result of spatial feedbacks between landscape structure and process. As the
 266 structure imparted by the vegetation changes, it leads to process feedbacks which lead to further
 267 changes in landscape structure. For example, the formation of mounds under shrubs leads to
 268 changes in infiltration and flow pathways, which become reinforced by increased erosion rates
 269 (Parsons *et al.*, 1996). Okin *et al.* (2009) provided a generic overview that suggested that
 270 connectivity across a range of processes, such as water and wind transport, animal activity and
 271 fire, could be used to understand ongoing degradation in the US Southwest. Based on this
 272 overview, it is hypothesized here that a connectivity-based model, when combined with local
 273 processes – as derived from empirical work developed from the Islands of Fertility model – will
 274 provide a way out of the present impasse, and allow modelling studies to investigate multi-scale

275 pattern formation in desert ecosystems. Specifically, a connectivity-based approach may
 276 provide a process-based understanding for the development of a model with advective and
 277 diffusive components, that overcomes the limitations discussed above. Furthermore, if such an
 278 approach rests upon a theoretical framework with a sound process-based understanding, it can
 279 thus be parameterized with data that can be measured and have a clear physical meaning.
 280 Testing the ability of such a model to reproduce the spatial and temporal progression of
 281 vegetation patterns, will enable it to be used with confidence to generate testable predictions
 282 relating to the function of dynamic desert ecosystems in response to a variety of endogenous and
 283 exogenous disturbances.

284

285

286 **3. Modelling Framework**

287 The ecogeomorphic modelling framework proposed here (figure 1) meets the characteristics
 288 outlined above. In this framework, the environment is represented as spatially related locations
 289 that may be inhabited by different vegetation types and quantities. Processes operating within
 290 the landscape are considered as being either vertical or lateral. Vertical processes act as inputs
 291 (e.g. rainfall, or atmospheric deposition of nitrogen) or outputs (e.g. losses of biomass through
 292 herbivory), while lateral processes act entirely within the system (e.g. movement of propagules
 293 by water). These processes provide locations with resources and propagules (R&P) that are the
 294 main drivers of plant growth and recruitment. Resources are defined in the model as abiotic
 295 materials that plants need to grow and propagate. Although the model can include any number
 296 of resources, emphasis has been placed on water and nitrogen as the principal limitations on

297 growth in deserts (Fitter and Hay, 1985; Whitford, 2002, p14; Wainwright, 2009). Propagules
 298 are biotic materials required for reproduction (e.g. seeds, tillers, cladodes).

299
 300 The lateral processes controlling the movement of R&P can be expressed as a set of vectors.
 301 Three vectors are included in the model: water, wind and animals. However, in principle any
 302 number could be included. The ability of vectors (in the sense of a geometric entity, having both
 303 a direction and magnitude) to redistribute R&P within the landscape is controlled by external
 304 factors. These external factors are termed “lateral externalities”. For instance, the movement of
 305 propagules by wind depends on wind speed and direction (Okin *et al.*, 2001). The movement of
 306 material by water depends on raindrop energy, the infiltration-excess runoff, surface topography
 307 and the characteristics of the material being moved (Parsons *et al.*, 2004). Animals move
 308 material according to foraging strategies. For instance, large grazing mammals can move large
 309 amounts of organic carbon and nitrogen in their guts over great distances, while propagules can
 310 be moved internally or externally as burrs and cladodes attached to the hide of the animal
 311 (Turchin, 2003).

312
 313 Similarly, the vertical processes that move R&P into and out of the landscape are controlled by
 314 external factors, termed “vertical externalities”, which include infiltration, leaching, evaporation
 315 and wet/dry deposition of nutrients or seeds. Direct disturbance of biomass is controlled by
 316 disturbance factors, termed “disturbance externalities”, which include, destruction or removal of
 317 biomass by fire, disease or herbivory.

318

319 The operation of the vectors which control the movement of R&P can be subdivided into two
 320 broad process states. These two process states are advection and diffusion. The important
 321 advective processes are concentrated overland flow (Wainwright *et al.*, 2008a), aeolian transport
 322 through large interplant gaps and the movement of, typically, large animals through the
 323 landscape. They are a function of lateral externalities such as wind strength, flow hydraulics, or
 324 the type and number of large animals. The important diffusive processes are splash, local
 325 distribution by small eddies and movements caused by small animals.

326

327 For the purpose of the model, connectivity quantifies the extent to which individual cells of the
 328 landscape may receive a subsidy as a result of the operation of a specific vector. The
 329 effectiveness of vectors to move R&P in the direction of flow depends on the spatial
 330 arrangement of what are termed connected pathways (Bartley *et al.*, 2006). For wind, connected
 331 pathways are quasi-linear, aligned with the direction of the wind and terminated when wind
 332 intersects a plant (Okin and Gillette, 2001; Okin *et al.*, 2009). For water, gross pathways follow
 333 the energy slope and net pathways follow the aspect. For animals, the definition of connected
 334 pathways depends on the behaviour of individual species, for example, for large grazing
 335 mammals; a connected pathway is defined as contiguous areas with palatable biomass above a
 336 threshold amount (Turchin, 2003; Thornes, 2007). In the model, all connected pathways
 337 terminate in sinks. For example, a shrub would terminate the wind vector, a vegetation patch
 338 would terminate the water vector, and an unvegetated area would terminate the grazing vector.
 339 A diagrammatic representation of connected pathways and their terminations is shown in figure
 340 2. The behaviour of the biomass controls the number and arrangement of connected pathways.
 341 However, because connectivity is also an emergent property of the model, reflecting the

342 feedbacks between the vectors and vegetation growth and death, it is also an independent
 343 measure of the ability of the model to represent the dynamics of desert ecosystems: if the
 344 processes in the model operate in a way that is compatible with reality, then connected pathways
 345 should be observable in real landscapes, and rates of change should also be equivalent.
 346

347 Not all distributions of R&P are amenable for movement by vectors. Labile nitrogen beneath
 348 canopies, for instance, cannot be moved by infiltration-excess runoff occurring in plant
 349 interspaces. The extent to which R&P are amenable to movement by vectors is specified in the
 350 model as “availability”. Availability is controlled by the biomass content in each spatial location.
 351 For wind and water, which cannot (under non-drought conditions) remove material from under-
 352 canopy areas, R&P may only be removed from plant interspaces. Animals, in turn, can forage
 353 only where there is a significant amount of palatable forage.
 354

355 Demographic processes (recruitment, establishment, growth and mortality) control the amount of
 356 biomass of individual species (e.g., Coffin and Lauenroth, 1990). The spatial patterning of
 357 structurally diverse vegetation controls the strength of vectors acting on the land surface (e.g.,
 358 Okin and Gillette, 2001). For example an open plant community would allow for a greater
 359 number of connected pathways along which wind and water could operate, but some species of
 360 plants do not provide palatable forage and so reduce the number of grazer-connected pathways.
 361 The latter reduction in connectivity in turn affects the redistribution of R&P, resulting in sinks of
 362 resource where plants are more likely to become established and survive. Therefore, this model
 363 explicitly considers relationships among the forces that control R&P movement via vertical and
 364 lateral externalities and vegetation distribution via connectivity. As a first approximation, it is

365 assumed that soil mechanical properties (including density and hardness) may be neglected and
 366 so plant growth is assumed to be equally possible in all locations (Bugmann and Solomon, 1995;
 367 Higgens *et al.*, 1996; Starfield, 1996), though, as with the number of vectors, such variables
 368 could, in principle, be incorporated into the model.

369

370

371 **4. Numerical Implementation of the Model**

372 In order to represent the spatial arrangement and structure of the simulated ecosystem, the area
 373 under consideration is divided into a grid of equally spaced nodal points enclosed by square cells
 374 of equal size. The number of cells used is specific to the particular implementation. The
 375 placement of these points coincides with the physical boundaries of the grid. A general point 'P'
 376 and its neighbours are identified using a Moore neighbourhood of the eight neighbouring points
 377 (figure 3). In order to limit the effects of numerical boundary conditions, cells are always
 378 wrapped (i.e. to form a cylinder) in the direction perpendicular to the vectors using a simple up-
 379 and-down procedure (Furukawa *et al.*, 2000). For water the direction of the vector is defined
 380 (arbitrarily and in this implementation of the model) as north to south implying a sink of water at
 381 the southern edge of the grid. Therefore, the line of action for grazers is south to north (that is,
 382 away from the implied water source, e.g. Lange *et al.*, 1984). For both water and grazers, cells
 383 are consequently wrapped across the east and west boundaries. Wind is allowed to operate in
 384 any direction across the grid, but in the present implementation, the line of action is east to west
 385 so the north and south boundaries are wrapped. A periodic boundary condition is applied in the
 386 direction aligned with the vector (Leach, 2001) when the grid represents a terrain with no slope.

387 This condition was also applied to the application of the model described below to demonstrate
 388 that numerical boundary conditions did not affect the model results.

389
 390 The model operates with an annual time step. For each cell the change in resource and
 391 propagules (represented through the term Z) in each time step (t) is as a result of three actions.
 392 These are the actions of vertical processes (Q_V), which add or remove material to the cell, the
 393 lateral processes (Q_H), which redistribute material in the grid, and the action of plant species (U),
 394 which varies depending on whether resources or propagules are being considered: where abiotic
 395 resources are considered, U represents a consumption term and where propagules are considered
 396 U represents a production term. The change in each (i) abiotic resource and biotic propagule for
 397 each cell can be expressed in differential form:

$$\frac{dZ_i}{dt} = \frac{dQ_{Vi}}{dt} + \frac{dQ_{Hi}}{dt} + \frac{dU_i}{dt} \quad (1)$$

398
 399
 400
 401 The model needs to be general enough to allow a range of physical processes to operate within
 402 the grid. However, different physical processes will operate over different spatial and temporal
 403 scales. The different spatial scales are accounted for by parameterizing the model according to
 404 the size of the cells used in the implementation, but a different procedure must be used for
 405 defining the various temporal scales. For example, a summer monsoon-type rainstorm will
 406 quickly lead to the generation of overland flow (Parsons *et al.*, 1997), whereas processes
 407 allowing this water to infiltrate – particularly to deeper soil layers – may take considerably
 408 longer (Hillel, 2004; Wainwright, 2009; Wainwright and Bracken, 2011). In order to calculate
 409 the physical responses of such a system, Peters (2002a) used daily totals (for example rainfall

410 rate), which were summed to monthly totals, whereas the more appropriate yearly values for
 411 biomass were aggregated over each month. Equation 2 is therefore cast to reflect different
 412 temporal scales (represented by θ , ϕ and ψ) summing to the same time scale (t).

413

$$414 \quad Z_i = \int_t^{t+1} \left[\left(\frac{dQ_{Vi}}{d\theta} \right) \frac{d\theta}{dt} + \left(\frac{dQ_{Hi}}{d\phi} \right) \frac{d\phi}{dt} + \left(\frac{dU_i}{d\psi} \right) \frac{d\psi}{dt} \right] \quad (2)$$

415

416 The different timescales imply that different processes are applied in a strict order (from fastest
 417 to slowest) in the calculation, and each process is enclosed within its own iterative loop. The
 418 numerical solution is therefore formulated so that the first calculation procedure adds or removes
 419 R&P by vertical processes. Subsequently the R&P are redistributed by lateral processes. Finally
 420 the biomass is allowed to respond to these new R&P distributions at the end of the time step
 421 (figure 4). All results presented below are for the aggregated effects of these timescales at an
 422 annual resolution.

423

424 **4.1 Vertical Processes**

425 Vertical processes are those that can move R&P into and out of the landscape without
 426 intermediate movement across the surface of the grid. To represent different processes, the Q_V
 427 term is divided into two parts. The first part ($Q_{V_{external}}$) represents those processes controlled
 428 wholly by vertical externalities, and thus operate independently of the biomass in the cell (e.g.
 429 precipitation rates); the second part ($Q_{V_{internal}}$) comprises those vertical processes linked to cell
 430 biomass (e.g. infiltration rates). For simplicity of implementation, the effects of disturbance
 431 externalities are also included within vertical processes and can be internal or external according

432 to whichever disturbance externality is being considered. The term Q_V is therefore a
 433 representation of the output of suitable sub-models to describe these phenomena:

434

$$435 \quad QV_i = QV_{i,external} + QV_{i,internal} \quad (3)$$

436

437 By defining a spatial grid of nodal points, the model has the flexibility to include appropriate
 438 sub-models for spatial variability that apply to the process in question.

439

440 **4.2 Lateral Processes**

441 The extent to which R&P are amenable for movement by vectors, is specified by the
 442 “availability” (A) term, which is itself a function of biomass in each spatial location. We define a
 443 maximum biomass (B_{max}) for each species that can exist in each cell and a linearly decreasing
 444 amount of R&P (Greene *et al.* 1994, Morgan 1996) is available to the water and wind vectors as
 445 the actual biomass in the cell increases as:

446

$$447 \quad A_{i,cell} = \left[Z_i - Z_i \frac{B}{B_{max}} \right]_{cell} + Q_{up,i} \quad (4)$$

448

449 while a linearly increasing amount of R&P is available to grazers as biomass increases:

450

451

$$452 \quad A_{i,cell} = \left[Z_i \frac{B}{B_{max}} \right]_{cell} + Q_{up,i} \quad (5)$$

453

454 For all three vectors, the availability of propagules increases with biomass (equations 4 and 5).

455 The availability term also includes the R&P moved into the cell from its up-vector neighbour

456 ($Q_{up,i}$).

457

458 When the connected pathways are terminated, the sum of the resource that has been entrained by

459 the vector along the connected pathway is redistributed from the cell terminating the connected

460 pathway, i.e. R&P movement becomes a diffusive, internally controlled process. The form of

461 this redistribution is modelled by a series of convolution matrices that are specific to each

462 transport vector and each sink (e.g. shrub or grass). It is assumed that upon encountering a cell

463 terminating a connected pathway, and in all subsequent vegetated cells along the line of action of

464 the vector, some proportion of the entrained R&P (α , β , γ , ε and ζ according to the relative

465 position: see equation 6) is redistributed to the eight cells that surround the disconnected cell.

466 First, the R&P moved down-flux is combined with the available R&P of the current cell into a

467 single term, Q_{ADV} , (the sum of all the available R&P). It is considered that part of this R&P

468 remains in the current cell, lateral distributions of R&P from the current cell are symmetrical, but

469 redistributions along the line of the vector can be asymmetrical. These descriptions are

470 summarised in equation set 6, using the water vector as an example. Except for the R&P that is

471 advected to the south cell, the R&P that is redistributed to the neighbour cells is unavailable for

472 further movement by the vectors in the current time step. For the other vectors, the equation set

473 is rotated relative to the appropriate direction of operation of the vector.

474

NW = $\frac{\epsilon}{2} Q_{ADV}$	N = βQ_{ADV}	NE = $\frac{\epsilon}{2} Q_{ADV}$
W = $\frac{\alpha}{2} Q_{ADV}$	Current Cell = $Q_{ADV} (1 - \alpha - \beta - \gamma - \epsilon - \zeta)$	E = $\frac{\alpha}{2} Q_{ADV}$
SW = $\frac{\zeta}{2} Q_{ADV}$	S = γQ_{ADV}	SE = $\frac{\zeta}{2} Q_{ADV}$

(6)

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This description links the amount of R&P that is moved to the magnitude of the external controls, in the sense that larger rainfall amounts result in a greater amount of R&P moved via the lateral processes compared with drought years. It should be noted that the length and spatial arrangement of the connected pathways depends upon biomass response, which itself is a function of lateral and vertical processes that occurred during previous time steps.

Only the R&P that is added in the current time step is moved in this way. Propagules are either established as seedlings or fail to establish before subsequent calculation steps, and any resource remaining from previous time steps is considered as unavailable to the lateral processes and moved down through the soil layers by the vertical processes. All resource, whether added and moved in the current time step or accumulated from previous time steps is then available to be used by the biomass for growth and propagation.

490 The derivation of the convolution matrices is analogous to the discretization of partial
 491 differential equations to describe the diffusion of R&P around the current cell, and is based on
 492 the finite volume method for computational fluid dynamics (e.g. Versteeg and Malalasekera,
 493 1995). As such, this model can be implemented (via equation 6) analytically, where explicit
 494 equations control spatio-temporal dynamics, or numerically, where dynamics are controlled by a
 495 combination of analytical equations and neighbourhood-based rules. In the present
 496 implementation, these convolution matrices are specified as neighbourhood-based rules that
 497 encapsulate the detailed biophysical processes that result in the deposition, and patterns of
 498 deposition, of material in the vicinity of plants. This neighbourhood-rule approach allows the
 499 transport of material by vectors to be included explicitly without the computational burden of
 500 having to model the sub-grid-scale physics explicitly.

501

502 **4.3 Biomass Response**

503 Any suitable model for vegetation-growth dynamics can be included through the term U (in
 504 equations 1 and 2). The present implementation of the model is designed to test the extent to
 505 which local redistribution of R&P can lead to emergent patterns of desert vegetation, and so to
 506 include a logistic growth equation here (e.g. Thornes and Brandt, 1993) would mask the effects
 507 of changing R&P when the biomass content of a cell is near to zero or the carrying capacity
 508 (Kot, 2001). Biomass is thus allowed to change linearly in response to new resource levels in the
 509 cell.

510

511 The change in biomass (B_j) for each species j is calculated using the sum of each resource in the
 512 cell (R_{TOT}), which includes the redistributed resource from the current time step plus any

513 resource remaining from previous time steps, which is stored in the lower soil layers. The actual
 514 change in biomass is computed from the most limiting of the resources.

515

$$516 \quad \Delta B_{j,i} = \left[\frac{R_{tot,i} - B_j M_{i,j}}{E_{i,j}} \right] \quad (7)$$

517

518 M is the requirement of each *i* resource to maintain 1 g of perennial material in the plant (used
 519 here to describe the resource requirement for maintenance of biomass), and E is the resource
 520 needed to yield 1 g of new leafy material (used here to describe the resource requirement for new
 521 growth) (Peters, 2002a). This method allows the model to describe vegetation response by
 522 process, as multiple species may be parameterized using data that can be measured in the field.
 523 As such, multiple actual species are represented, unlike the majority of previous spatial
 524 modelling work where generic grass- or shrub-type species have been considered as broad
 525 functional types (Thornes and Brandt, 1993; Koppel and Rietkerk, 2004; Istanbuluoglu and
 526 Bras, 2006; Klausmeier, 1999; HilleRisLambers *et al.*, 2001; Rietkirk *et al.*, 2002, Coueron and
 527 Lejeune, 2001; Barbier *et al.*, 2006).

528

529

530 **5. Model Implementation**

531 In order to test the numerical model, it has been implemented with respect to parameters and
 532 conditions in the deserts of the south-western USA. Two sets of simulations have been
 533 conducted, but only the latter are presented here. In the first, a series of simulations was carried
 534 out to explore the behaviour of the model. These simulations are reported in the Appendix.

535 They showed agreement of the behaviour of our model with observed characteristics of desert
 536 vegetation, giving us confidence to use the model test hypotheses of vegetation change in the
 537 American Southwest.

538
 539 The American Southwest is one of the many regions in the world where invasion of woody
 540 shrubs into desert grassland has been observed (Schlesinger *et al.*, 1990; Archer *et al.*, 1995;
 541 Allred, 1996). Increasing aridity (e.g. Archer, 1995; d’Herbes *et al.*, 2001) and overgrazing (e.g.
 542 Archer, 1995; Okin *et al.*, 2001; Nash *et al.*, 2004; Westoby *et al.*, 1989; Gibbens and Lenz,
 543 2005) are commonly used explanations for shrub invasions into grasslands. Thus, here we use
 544 the model to investigate the viability of these explanations both individually and in combination,
 545 and to propose testable hypotheses of why changes in desert ecosystems are initiated. A further
 546 advantage of the use of this region for an implementation of the model is that it is relatively data-
 547 rich, and thus some information exists that can be used to provide an independent test of the
 548 model output.

549
 550 The specific site chosen for model implementation is the Jornada Basin Long Term Ecological
 551 Research site in southern New Mexico (32°37’ N, 106°40’ W, 1260 m a.s.l.). The Jornada Basin
 552 LTER was established with the task of quantifying the processes that have caused dramatic
 553 changes of structure and functioning of Chihuahuan desert ecosystems such as have been noted
 554 over the past 150 years, and links into pre-existing and continuing datasets collected by the
 555 USDA-ARS Jornada Experimental Range field station. An impressive array of data is available
 556 for the Jornada Basin, which makes it possible to look for temporal trends, spatial patterns and
 557 ecosystem changes over the 20th century (Havstad *et al.*, 2006; Jornada LTER, 2010;

558 Wainwright, 2005; Yao *et al.*, 2006). Consequently, the region, and particularly the LTER site,
 559 is rich in data with which to parameterize our model. The Jornada Basin LTER, experiences
 560 severe drought (Palmer Drought Index between -3 and -4) (Nicholson, 1979; 1981) every 20 to
 561 25 years, and extreme drought (Palmer Drought Index of <-4) occurs every 50 to 60 years.
 562 Livestock were introduced from Mexico during the early part of the 16th century, but grazing was
 563 limited in the Jornada Basin owing to the lack of surface water until the sinking of the first wells
 564 in 1867 following the Homestead Act of 1862. Since then it has supported a mainly beef
 565 rangeland livestock industry (Gibbens *et al.*, 2005; Havstad *et al.*, 2006). For many arid and
 566 semi-arid ecosystems the amount of biomass supported per unit area of primary production is
 567 approximately an order of magnitude greater under rangeland livestock production than under
 568 natural non-agricultural conditions. For example, in the Jornada the biomass of native
 569 consumers is approx 0.03 g m^{-2} , which consume less than 5 g m^{-2} of forage per annum compared
 570 to a conservative stocking rate of cattle of 1.7 g m^{-2} , which consume $8\text{-}14 \text{ g m}^{-2}$ per annum on the
 571 same grassland (Havstad *et al.*, 2006 but see also Pieper *et al.*, 1983).

572

573

574 **5.1 Model Parameterization**

575 A realistic test of the conceptual model should be undertaken with reference to specific localities
 576 and specific species if insights beyond broad generalizations are to be gained. Such site-specific
 577 insights are a prerequisite for informed management interventions (Westoby, 1980).

578 Accordingly, parameterization is based, as far as possible on data obtained from field
 579 measurements at the Jornada LTER, or elsewhere in the US Southwest. For the implementation,

580 we use a planar 50 × 50 m grid that is subdivided into cells of 1 m² with a downslope gradient of
 581 2° north-south. A summary of the model input conditions is presented in Table 1.

582

583 **5.1.1 Biomass**

584 The encroachment of grassland by woody shrubs may involve several species, but in this
 585 implementation of our model, plant demographic processes were parameterised using two
 586 species that are indicative of the grass to shrubland transitions observed in the deserts of the
 587 south-western USA (Humphrey and Mehrhoff, 1958; Schlesinger *et al.*, 1990; Peters, 2002a;
 588 Peters, 2002b). These two species are *Bouteloua eriopoda* (Torr.) Torr. (black grama) and
 589 *Larrea tridentata* (DC.) Coville (creosotebush). The first represents a typical desert grass, which
 590 is the dominant species in many hot desert grasslands of the southwest USA (Smith *et al.*, 1996;
 591 Nelson, 1934). Black grama typically occurs on rocky or sandy mesas and open ground, with
 592 well drained sandy and gravelly soils (Humphrey, 1958) and is particularly abundant in the
 593 Chihuahuan desert (Yao *et al.*, 2006; Peters, 2002b).

594

595 Black grama often shows an association in upland areas with our second simulated species,
 596 creosotebush (Gardener, 1950). Creosotebush is a drought-tolerant, evergreen shrub and a
 597 dominant or co-dominant member of many plant communities in the Southwestern deserts
 598 (Humphrey, 1958). It usually occurs in open, sparsely populated areas, but also appears as a
 599 transitional species in desert grasslands (Humphrey and Mehrhoff, 1958), and is noted to grow
 600 on bajadas, gentle slopes, valley floors, sand dunes and in arroyos, typically on calcareous, sandy
 601 and alluvial soils that are often underlain by a caliche hardpan (Went and Westergaard, 1949).
 602 Creosotebush occurs as far south as north-central Mexico, and as far north as central Nevada

603 (Ackerman and Bamberg, 1974; Pockman and Sperry, 1997), where average annual rainfall
 604 ranges from 100 to 300 mm (Castellanos and Molina, 1990).

605

606 **5.1.2 Externalities and Vertical Processes**

607 The instrumental record for rainfall at the Jornada begins in 1914, so longer term data for rainfall
 608 are only available by reconstructions. Tree-ring data have been used to reconstruct climate data
 609 in the US Southwest for the time period extending over the last few hundred years (d'Arigo and
 610 Jacoby, 1992), that includes the period of introduced cattle grazing. Data are available from the
 611 International Tree Ring Databank for three locations within a 50-km radius of the Jornada Basin,
 612 which were used by Wainwright (2005) to reconstruct a common sequence of the Jornada
 613 climate (figure 5a) extending from 1659 to 1970. Comparison with the more recent instrumental
 614 record (approximately 80 years long) suggests that although the retrodictions can capture
 615 extreme events, they tend to underestimate the magnitudes of these events (figure 5b). Cycles
 616 were present similar to those in the instrumental record, and drought conditions were retrodicted
 617 for the years 1676, 1790, 1721, 1723, 1736, 1872 and 1912. Greater inter-annual variability (in
 618 terms of number of rain days, rainfall totals and annual moisture balance) were noted until the
 619 late 18th century, whereas the 20th century seems to be particularly anomalous with long wet
 620 periods alternating with dry spells. The reconstructed rainfall record provides a regional input in
 621 terms of precipitation to our model, and includes disturbances due to drought implicitly.

622

623 Cattle-stocking levels over the 20th century for which historical data are available (Havstad *et al.*,
 624 2006) are shown in figure 6. Disturbance due to grazers is modelled explicitly by simulating the
 625 removal of a specified quantity of palatable biomass from each vegetated cell, in each year. A

626 summary of the simulated grazing levels is presented in table 2 which reports the minimum and
 627 maximum of harvest rates for cattle under three different stocking levels described as
 628 conservative, recommended and overgrazed by Havstad *et al.*, 2006. In our simulations, we used
 629 the central value of the reported consumption range.

630
 631 **5.1.3 Lateral Processes**

632 In this implementation, three vectors are defined: water, wind and animals. As the action of
 633 grazers is being modelled as a disturbance externality, the animal vector in this implementation
 634 represents the action of cattle. The model requires rules for the proportion of resources and
 635 propagules that can be moved by each vector, and, although all the parameters used in this model
 636 could be measured at the field site, as yet much of this information is not available in a form that
 637 can be implemented numerically. Consequently, we have inferred the R&P availability and
 638 convolution matrices from experimental work which alludes to the mechanisms by which vectors
 639 might operate (Parsons *et al.*, 1992; Parsons *et al.*, 1997; Parsons *et al.*, 2004; Parsons *et al.*,
 640 2006a-b; Okin and Gillette, 2001; Okin *et al.*, 2001; Okin *et al.*, 2009), and also from process-
 641 based modelling studies (Scoging *et al.*, 1992; Parsons *et al.*, 1997; Wainwright *et al.*, 1999;
 642 Wainwright *et al.*, 2002; Wainwright *et al.*, 2008a-c). We specify that, of the resource added to
 643 each cell per time step by the vertical processes, water cannot be moved from cell to cell under
 644 the action of wind or grazers, and we allocate 45% of the nitrogen to be amenable to movement
 645 by water, 45% to be amenable to movement by wind and 10% to be amenable to movement by
 646 grazers. These proportions are arbitrary, but based upon the argument that wind and water have
 647 equal access to nitrogen in the soil, whereas grazers have access only via vegetation uptake.
 648 The redistribution of R&P from vegetated cells is effected according to the species-specific rules

649 defined in table 3a for water-disconnected locations, table 3b for wind-disconnected locations
 650 and table 3c for grazer-disconnected locations.

651

652 **5.1.4 Resources and Propagules**

653 Aside from rainfall and grazing externalities, model parameterization requires information about
 654 the abiotic resources necessary to support plant activities. Although the model can be
 655 parameterized to accommodate any number of resources, we have focused on two in this
 656 implementation: water and nitrogen. The water input to the model is provided through the
 657 descriptions of rainfall described previously.

658

659 It is, however, difficult to find suitable parameterization data for nitrogen. The analysis of the
 660 plant-available nitrogen in the soil is not a particularly useful measure of the total nitrogen
 661 available to plants, since nitrogen released by microbes can be rapidly taken up by plants and
 662 never appear in the soil pool (Gallardo and Schlesinger, 1992). This issue is noted to be
 663 particularly important in semi-arid environments (Clark and Tilman, 2008). The point at which
 664 nitrogen becomes limiting to plant growth at the Jornada, under conditions when water is not
 665 limiting, must be answered in the short-term, as a longer-term perspective would allow the
 666 invasion of non-desert vegetation, with much higher productivity, assuming that such vegetation
 667 possesses the necessary adaptations to semi-arid environments. In the short-term, it seems that
 668 nitrogen is limiting whenever water (the dominant control in its own right and through its partial
 669 control on nitrogen availability) is non-limiting, so the present rate of nitrogen mineralization of
 670 up to $5 \text{ g m}^{-2} \text{ a}^{-1}$ (Loreau *et al.*, 2002) would be the lower limit to plant growth in various
 671 habitats. The rate of annual nitrogen mineralization has not been measured in much detail since

672 many researchers believe periods of rapid mineralization probably occur during a few wet
 673 months, and then no mineralization takes place for the rest of the year. Thus, nitrogen inputs
 674 were simulated as a constant values in this work (Baez *et al.*, 2007), but the amount of nitrogen
 675 that is used by the plants to support existing biomass, and for new growth was allowed to vary
 676 linearly with rainfall (Burke *et al.*, 1990; Peters, 2006a; Wainwright, 2009).

677
 678 After a resource has been added by vertical processes, and moved by the lateral processes, it is
 679 then used to support plant growth. Descriptions of resource use by the biomass (table 4) are
 680 based on data given in Peters (2002a), Maneta *et al.* (2008) and Wainwright (2009).

681
 682 The consumption of resource by the biomass is calculated using the sum of each resource in the
 683 cell, which includes the resource from the current time step plus any resource remaining from
 684 previous time steps, which are stored in the middle and deep soil layers. Three soil layers are
 685 simulated:

- 686 • The top layer, from depths of 0 to 100 mm. This layer contains resource added in the
 687 current time step, and subsequently redistributed by the vectors. The depth was selected
 688 to coincide with measured wetting-front depths (Martinez-Meza and Whitford, 1996;
 689 Wainwright *et al.*, 2008b-c; Parsons *et al.*, 1997).
- 690 • The middle layer, from depths of 100 to 350 mm. This layer corresponds to the
 691 maximum observed depth for root-channelized water in creosotebush (Martinez-Meza
 692 and Whitford, 1996; Scott *et al.*, 2008) and encloses the point of maximum root density
 693 for both creosotebush and black grama (Martinez-Meza and Whitford, 1996; Sun *et al.*,
 694 1998; Gibbens and Lenz, 2001; Peters, 2002a).

- 695 • The deep layer, from depths of 350 to 1500 mm. This layer corresponds to the maximum
696 observed depth of creosotebush roots (Peters, 2002a; Gibbens and Lenz, 2005).

697

698 Black grama is an intensive exploiter of water and derives the majority of its moisture through
699 dense root networks in shallow soil layers that enable it to exhibit rapid growth and water
700 absorption following rainfall. Black grama tolerates short droughts, and recovers rapidly from
701 stress when water is available (Burgess, 1995). These factors are reflected in the higher
702 proportion of roots in our simulated top layer, which allows grass to access the resources that are
703 added in each time step ahead of creosotebush. During conditions with adequate water input, the
704 top and middle layers will receive a greater quantity of water, and this will favour the grass
705 (Walter, 1971; Thornes, 1990). Creosotebush has a greater proportion of roots in the middle and
706 particularly the deep layers where unused resource from previous time steps is stored. During
707 times of resource stress, creosotebush can access this store ahead of black grama, reflecting the
708 greater drought tolerance of shrubs (Walter, 1971; Herbel *et al.*, 1972; Casper and Jackson,
709 1997). These arguments represent competition processes between the species (Thornes, 1990).

710

711 The proportion of resource in each layer that can be used by each species is a function of root
712 biomass (table 5, figure 7). The above-ground biomass is converted to below-ground biomass
713 using the relationship proposed by Peters (2002a) wherein black grama root biomass is estimated
714 to be 1.44 times greater than the above-ground biomass, and the root biomass of creosotebush is
715 estimated to be equal to the above-ground biomass. In order to distribute root biomass among
716 the layers, we follow the method described by Peters (2002a) which was itself based on the
717 analyses of root distributions of a large number of grassland species in the US by Sun *et al.*

718 (1998). The method assumes that root biomass increases linearly to a species-specific depth,
 719 then decreases allometrically to the maximum depth. Parameterization data were all obtained
 720 from Peters (2002a).

721
 722 If creosotebush is the dominant species in the cell, then a proportion of the top layer resource,
 723 (equal to the proportion of creosotebush biomass in the cell), is directly channelled into the
 724 middle and deep layers (Martinez-Meza and Whitford, 1996; Abrahams *et al.*, 2003). The
 725 movement of resource through the soil layers acts as a facilitation term. Once established,
 726 creosotebush is thought to improve sites for the annual plants that grow beneath its canopy by
 727 trapping sediment, organic matter and propagules, and by increasing water infiltration and
 728 storage (Bainbridge and Virginia, 1990). Although this description of channelization should
 729 strictly apply only to water, nitrogen is also handled in the same way in the model
 730 implementation to reflect the ability of a plant with access to deep water to use more of other
 731 nutrients (Martinez-Meza and Whitford, 1996), and it is reasonable to assume that the water
 732 contains significant amounts of dissolved nitrogen (Schlesinger *et al.*, 1990; Schlesinger and
 733 Peterjohn, 1991; Grimm and Railsback, 2005; Brazier *et al.*, 2007; Michaelides *et al.*, 2012;
 734 Turnbull *et al.*, 2010).

735
 736 In each cell, if there is insufficient resource at a particular time step to satisfy maintenance
 737 requirements, the biomass is reduced. When the outcome is a loss of biomass, and this loss is due
 738 to insufficient water, the model allows all water to be used but no other resources are consumed
 739 (Hooper and Johnson, 1999). When the loss of biomass is due to a deficit of nitrogen, the model
 740 allows the biomass to consume all resources in sufficient amounts to maintain (as far as possible)

741 existing biomass. Under conditions of biomass loss, no propagules are generated. These rules
 742 reflect some of the observed adaptations of desert vegetation to survive extremes of climate
 743 (Walter, 1971; McClaran and Van Devender, 1995).

744

745 Descriptions of propagule movement are also inferred from literature. Although black grama
 746 provides excellent forage, populations are damaged by grazing as these plants rely heavily on
 747 stoloniferous regeneration (Gosz and Gosz, 1996; Canfield, 1948). Whilst these means of
 748 reproduction are effective under arid conditions, they do not promote extensive migration.
 749 Consequently, black grama is slow to colonize adjacent areas (Brown and Gersmehl, 1985).
 750 This effect is compounded by the low viability of its rarely produced seeds (Nielson, 1986). The
 751 effect of droughts and grazing are to decrease tuft area, which allows for greater wind erosion of
 752 the upper loose soil litter layer required for stolon rooting. Creosotebush is a stable member of
 753 desert plant communities owing to its primarily vegetative method of reproduction via cloning
 754 (Romney *et al.*, 1989; Cody, 1986). Germination of seeds is rare, and the rate is reported to be
 755 less than 20% outside of the optimal summer rainfall of between 75 and 150 mm (Ackerman and
 756 Bamberg, 1974). The seeds are primarily adapted for tumbling, as they are too heavy for lofting
 757 and the trichomes are not stiff enough to penetrate animal skin therefore not adapted to animal
 758 dispersal (Chew and Chew, 1970).

759

760 In the model, we simplify these complex conditions of propagule production and movement by
 761 allowing propagules to be generated whenever a positive growth rate is recorded. For our
 762 purposes, propagules are a species-specific proportion of the new growth of each plant, and a
 763 proportion of these propagules is available to the vectors for redistribution within the

764 environment. A small proportion of this annual new growth is allowed to move under the action
 765 of the vectors in the next time step, and will become established in new cells only if resource
 766 levels in these new locations are sufficient to support an increase in biomass. The majority of the
 767 propagules for each species will move to adjacent cells by diffusion to represent the
 768 predominantly asexual method of reproduction utilised by desert plants. If some propagules are
 769 moved to a connected cell, they are then dispersed along connected pathways by the vectors
 770 (Barbour, 1969; Miller and Donart, 1979).

771
 772 For both species, water availability is the primary controlling factor in terms of propagule
 773 dispersal (Aguiar and Sala, 1999). This control is modelled by having the majority of propagules
 774 following the line of action of the water vector (i.e. downslope). These parameter values can be
 775 changed to account for the different germination probabilities of individual species, but for the
 776 simulations presented here, the values are fixed (table 6). Species with seeds easily transported
 777 by the wind could be similarly moved in the model along the wind direction.

778
 779

780 **5.2 Descriptions of the simulations**

781 The simulations have been carried out to test hypotheses of the different rôles of precipitation
 782 and grazing in explaining woody shrub invasion and of the causes of spatial variability in
 783 response to drought. Conflicting results exist in the literature evaluating the effects of temporal
 784 variability in precipitation on vegetation. It is suggested that one cause of this conflict may be
 785 the consequence of the representation of the rainfall pattern in a model. To assess the extent to
 786 which this is the case, four simulations have been undertaken in order to explore different levels

787 of complexity in the representation of rainfall on the resultant vegetation patterns and their
 788 interactions with grazing pressure. Simulations were characterized by (a) stochastic rainfall with
 789 no temporal autocorrelation; (b) rainfall reconstructed from the tree-ring record for the period
 790 1659-1970; (c) as (b) but with variable grazing levels. For simulation (a) the stochastic rainfall is
 791 generated from the mean and standard deviation of the reconstructed rainfall of simulations (b)
 792 and (c). Finally, in simulation (d) we use the model in conjunction with the 80-year
 793 instrumented rainfall-data record to examine reported differences in response to the same
 794 climatic conditions.

795

796 **5.3 Initial conditions**

797 All simulations were initialized from the same randomly generated landscape (Table 1) which
 798 included a random distribution of black grama biomass and a uniform distribution of shrub
 799 biomass (to represent a seedbank). An initial biomass of 60 g m^{-2} was specified for black grama.
 800 This initial value was then perturbed by a low level of white noise (a random signal with a flat
 801 power spectral density, in this case, by generating a pseudorandom matrix of numbers lying
 802 between 0 and 1 with an average value of 0.5). This procedure follows the method of Coueron
 803 and Lejeune (2001). This method of perturbation of the biomass yields an initially random
 804 distribution of grass in each cell and the same initial random distribution was used as a starting
 805 point in all of the simulations presented here. The initial biomass of the shrubs was specified as
 806 10 g m^{-2} in all cells, to represent a seed bank. The randomly generated map of biomass is
 807 depicted in figure 8. The initial resource level in the mid and deep layers was set to 25 g/m^2 for
 808 water and 0.25 g/m^2 for nitrogen. Resource levels of the top layer were provided by the input of
 809 water and nitrogen in each time step.

810

811 In all of the following simulations, all three vectors (water, wind and grazers) operate to move
 812 R&P through the landscape. The effect of herbivory by grazers is included only in simulations c
 813 and d.

814

815 **5.4 Presentation of results**

816 We present the results of the simulations in two formats. In the first, we present three graphs
 817 showing a) the average change in grass and shrub biomass in the cells along a transect along the
 818 centre line of the grid; b) the average change in water and nitrogen in mid and deep soil layers
 819 for the same cells; and c) changes in connectivity for these cells. In the second format we
 820 display maps of grass and shrub biomass in each cell at selected times during the simulations.

821

822

823 **6. Results**

824 **Simulation a – Stochastic rainfall**

825 Results from Istanbuloglu and Bras (2006) have suggested that increased variability in rainfall
 826 and lower rainfall levels are mechanisms that on their own, and in combination with each other,
 827 will decrease the average grass biomass cover. This suggestion may be linked to the observation
 828 of Thornes and Brandt (1993), that more frequent woody plant encroachment and desertification
 829 are more likely to occur when the grass is in a degraded condition. Williams and Albertson
 830 (2006) argued similarly that some account must be taken of rainfall structure in models of
 831 dryland vegetation in order to understand the changes in a more meaningful way, (although they
 832 did not pursue this argument in their paper). For our first examination of the effects of rainfall

833 representation on shrub invasion, we have reproduced Williams and Albertson's model, which
 834 controls the statistical structure of annual rainfall by generating a synthetic rainfall time (t) series
 835 (P) that is represented by the equation

836

$$837 \quad P_t = \langle P \rangle + U_t + A_p \sin \left(\frac{2\pi t}{T_p} \right) \quad (7)$$

838

839 where U_t is an uncorrelated, log-normally distributed random variable with variance $\gamma\sigma_p^2$ where γ
 840 is a parameter lying between 0 and 1 that controls the partitioning of the total variance (σ_p^2)
 841 between uncorrelated (white) noise and correlated (sinusoidal) components (shown in figure 9a).
 842 A_p is the sinusoidal amplitude (mm) and T_p is the period (years). The mean annual rainfall $\langle P \rangle$
 843 is the long term average calculated from the tree-ring rainfall record as 228 mm a⁻¹ and the
 844 interannual variability of rainfall is represented by the coefficient of variance of rainfall [CV(P)],
 845 which is calculated as 49.0 %. The synthetic rainfall fluctuates in values between years and has
 846 no periodicity within the rainfall structure (figure 9b). The simulation was run for the same
 847 length of time as the length of the reconstructed rainfall record.

848

849 The response of biomass in our model correspondingly shows wild fluctuations in values.
 850 Depending on the features of the synthetic rainfall series, either of the two species is equally
 851 likely to become the dominant biomass in the grid without any bias towards the grass or the
 852 shrub. In the realization of the stochastic model shown in figure 10a, the grass and shrub
 853 continually alternate as the dominant species, and neither species shows any evidence of spatial
 854 reorganization.

855

856 When the grass is the dominant species in a cell, little available resource migrates to the deeper
 857 soil layers (figure 10b). However, when the shrub is the dominant species, a large amount of
 858 water resource is channelled to the deep layers. Transitions between one dominant species and
 859 another are accompanied by a change in connectivity (figure 10c). When a drought occurs, the
 860 biomass of both grass and shrubs is reduced, and when rainfall subsequently increases the grass
 861 and the shrubs both increase their biomass. The principal difference between the two plant
 862 species during recovery is growth rate. As the grass has a higher growth rate, it can recover
 863 slightly more quickly than the shrubs and suppress shrub invasion. In this sense, the variability
 864 in the rainfall itself inhibits shrub invasion, and confirms that the periodicity in rainfall is an
 865 important control on vegetation response.

866

867 **Simulation b – Reconstructed rainfall**

868 For this simulation the actual 312-year rainfall series that was reconstructed from tree-ring data
 869 was used. Results of same centre-line averages as in Figure 10 are presented in figure 11. The
 870 initial decline of grass biomass is caused by initiating the calculation with uniform resources in
 871 the mid and deep soil layers. Over the first 25 years of the simulation, the biomass adjusts to
 872 these arbitrary resource levels, and so these first few data points are excluded from further
 873 analysis. Figure 11a shows that after this initial period, although the average value of grass
 874 fluctuates, the shrubs are unable to become established.

875

876 It is noteworthy is that even after the most extreme drought (that of the 1950s), the grass
 877 population is able to recover and the shrub population continues to be suppressed. It is not
 878 unreasonable to expect that the same recovery of grass should be observed if droughts of a

879 similar magnitude occurred at earlier points in the simulation (c.f. McClaren and Van Devender,
 880 1995), assuming that the reconstructed data underestimate the magnitude of earlier droughts (see
 881 section 5.1.2).

882
 883 The middle layer water resource fluctuates markedly (figure 11b). As the average grass biomass
 884 increases, water levels decrease and *vice versa*. In spite of the accumulation of the water in this
 885 layer, the shrubs are not able to invade, which, assuming our model is faithfully reflecting the
 886 impacts of rainfall variability, suggests that some other mechanism apart from drought must be
 887 important in shrub-invasion processes.

888
 889 Connectivity (figure 11c) also fluctuates during the simulation. The average connectivity values
 890 show that whereas the magnitude of change in the grass population reflects the magnitude of
 891 change in the rainfall record, the response of the grass population tends to lag slightly behind
 892 changes in the rainfall (typically by two years). The lag is partly due to the unused resource
 893 being moved down through soil layers, and partly due to the structure of the rainfall in the sense
 894 that multiple wet years will increase patch biomass, which delays the effect of subsequent dry
 895 years on population decline (and *vice versa*). Although this delay may in part be an artefact of
 896 the calculation scheme, it is not inconsistent with the observations of the actual behaviour of the
 897 system as evaluated by Reynolds *et al.* (1999). The maximum number of connected cells also
 898 changes, related to a change in the spatial organization of the grass patterns. Therefore, the
 899 implication is that the temporal structure of rainfall plays a significant role in the spatial
 900 organization of vegetation, as well as its dynamic response.

901

902 In order to examine the changing spatial distributions of the biomass, maps of the distributions of
 903 grasses and shrubs are shown in figure 12. The biomass of grass decreases during times of water
 904 scarcity in a consistent and predictable way. During dry years, grass is first lost from cells
 905 containing the lowest biomass. These cells connect to the wind and water vectors and resource
 906 ‘flows’ into the next vegetated (i.e. disconnected) cell. This flow has the effect of concentrating
 907 resource into distinct spatial locations and allowing a higher biomass of grass to survive than
 908 would otherwise be possible were the resources more homogenously distributed (Humphrey,
 909 1958; Buffington and Herbel, 1965; Allred, 1996; Couteron and Lejeune, 2001; Barbier *et al.*,
 910 2006). At first, the vegetation loss occurs only in the lowest biomass cells. During prolonged
 911 periods of water stress, vegetation is lost from the downslope edge of the vegetation patch,
 912 because the water input provided by the vector is exhausted before it reaches this edge and
 913 consequently a ‘banded’ pattern is formed (e.g. noted in years 1783 to 1883). As noted in the
 914 Appendix, this banded pattern that is widely reported in the literature on desert vegetation. Over
 915 time, these bands become more fragmented. When the water input is increased, grass recovery
 916 initiates from all surviving grass cells. The relationship between the number of connected cells
 917 and the width of the grass band is a function of both the resource input and the biomass of the
 918 band.

919
 920 Meanwhile, the shrub biomass declines (as suggested by Goldberg and Turner, 1986) and by the
 921 time grass reaches its quasi-average value the initial shrub biomass has been reduced to an
 922 average value of almost zero. Where shrubs are able to survive, they do so only on the edge of a
 923 grass patch which has accumulated excess resources, and in effect, the grass patch acts as a nurse
 924 plant to the new creosotebush (McAuliffe, 1988). This quasi-static equilibrium level for the grass

925 controls the resource and propagule movement, by which it is meant that the pattern of resource
 926 movement in this simulation is predominantly lateral, locally limited, and observed in the top
 927 layers on the same spatial scale as the individual grass plants (Müller *et al.*, 2008). This result is
 928 consistent with evidence that the dominant species redistributes resource to suit its own
 929 colonization strategy, which has been noted by Westoby *et al.* (1989).

930
 931 The grass never quite reaches a stable equilibrium, irrespective of the duration of the simulation.
 932 Model runs of 1000 years were also carried out using both repetitions of the stochastic rainfall
 933 record and repeated cycles of the tree-ring record. In these model runs the grass population did
 934 not get any closer to reaching a steady equilibrium, which allows the idea that equilibrium is
 935 asymptotically reached, given an infinitely long time, to be rejected.

936
 937 **Simulation c – Reconstructed rainfall and variable grazing levels**

938 Simulations 2a and 2b indicated that the temporal structure of the rainfall is a causal factor
 939 leading to the generation of patchy vegetation, but the results also suggest that historical climatic
 940 conditions appear to be insufficient to cause the invasion of shrubs into grasslands. The
 941 introduction of grazers to the Jornada has often been cited as a reason behind woody plant
 942 encroachment (e.g. Archer, 1995) and so in this simulation, the variable rainfall input is
 943 combined with three different grazing intensities. These grazing levels are modelled as a
 944 disturbance externality by allowing some of the grass in each cell (as a percentage of the
 945 maximum cell biomass) to be removed. This percentage corresponds to the mid-point of three
 946 grazing intensities reported for the Jornada (table 2).

947

948 Figure 13a(i) show that with a conservative grazing intensity, the average biomass, resources and
 949 connectivity are little changed compared with simulation b (mean average grass and shrub
 950 biomass 30.6 g m^{-2} and 0.4 g m^{-2} , and 27.5 g m^{-2} and 0.6 g m^{-2} , respectively) where the effects of
 951 herbivory were not simulated (although the grazing vector did operate to move R&P in
 952 simulation b). Although average grass biomass levels (figure 13b(i)) are reduced in the
 953 simulation with “recommended” (Havstad *et al.*, 2006) compared with the “conservative”
 954 grazing intensities, there is little difference in the average shrub levels (mean average grass and
 955 shrub biomass 27.5 g m^{-2} and 0.6 g m^{-2} , and 27.8 g m^{-2} and 0.6 g m^{-2} , respectively). The
 956 connectivity values (figure 13(iii)) show that under recommended grazing levels, the effects of
 957 droughts become more pronounced with a greater number of cells connecting to the wind and
 958 water vectors, and the average connectivity suggests different spatial patterns in the vegetation
 959 occurs in response to the elevated grazing levels(maximum continuous connected cells 38 and
 960 38, respectively, and average connected cells 4.6 and 5.6, respectively). The “overgrazed”
 961 simulation (figure 13c) produces conditions that result in a dramatically reduced grass biomass
 962 and a much high higher average shrub biomass (18.7 g m^{-2} and 10.6 g m^{-2} , respectively). The
 963 effect of the increased shrubs is also seen in the average resource levels (figure 13c(ii)) where
 964 greater levels of water resources are able to penetrate the deep soil layers (different spatial scale
 965 arguments). By the time that the drought of the 1950s occurs, the biomass of the shrubs exceeds
 966 that of the grass (Schlesinger and Pilmanis, 1998). In spite of this change, the connectivity
 967 values (figure 13c(iii)) imply a very static pattern in the spatial pattern of biomass.

968

969 In order to interpret the patterns of connectivity in the overgrazed case (figure 13c(iii)), the
 970 spatial patterns of biomass for the overgrazed case are shown in figure 14. It can be seen that the

971 typical banded pattern is established early in the simulation. The resources that are moved by the
 972 vectors should be able to sustain a high level of grass in the disconnected cells (as in simulation
 973 b); however, the grazers remove some of this biomass. What would have been adequate resource
 974 becomes an excess resource on these grass patches, and the shrubs are able to colonise these
 975 areas of resource excess. Once the shrubs have become established in the locations shown in
 976 figure 14, the model identifies them as the dominant species, which has two consequences. First,
 977 the diffusion descriptions (table 3) pertaining to shrubs are used in place of the diffusion
 978 descriptions for grass. Secondly the shrubs are allowed to channel some of the water input to
 979 their cells directly to the middle and deep soil layers. At this point, the resource redistribution
 980 changes from a predominantly lateral process (where resources move to adjacent cells in the top
 981 soil layer) to a more vertical process where the shrubs are able to channel a higher proportion of
 982 water resources directly into deeper layers. During times of resource scarcity, these shrub
 983 patches contract and shrub biomass is reduced, but they are to some extent buffered from the
 984 effects of water shortage by the deep-water store (Thornes and Brandt, 1993). Thus the shrub
 985 community is quite stable, and therefore the spatial distribution of biomass and connectivity
 986 values also become more static. In the earlier simulations, the grass community was able to
 987 adapt to resource scarcity by expanding and contracting in patches as a function of connectivity.
 988 In the present case, the shrub population interrupts this connectivity and the grass patches
 989 become increasingly fragmented. At the end of the simulation, most of the grass survives only at
 990 the edge of the shrub patch (as observed in the field by McAuliffe, 1988).

991

992 **Simulation d – Instrumented rainfall data and a conservative grazing level**

993 Reynolds *et al.* (1999) and Yao *et al.* (2006), reported that different sites within the Jornada
 994 Basin have responded differently to the same climatic conditions: some stands of perennial grass
 995 became extinct before the drought of the 1950s, some during the drought, some immediately
 996 afterwards, and some not at all. Yao *et al.* (2006) used long-term cover data over a period from
 997 1915 to 2001 to identify this spatial variation in grass cover in the Jornada and, in the absence of
 998 a consistent causal factor, hypothesised that local transport processes for resources and
 999 propagules between patches must be somehow be important. In this simulation we use 80 years
 1000 of instrumented rainfall data to generate simulated grass responses in order to explore the
 1001 model's ability to test Yao *et al.*'s hypothesized explanation. Unlike previous simulations, the
 1002 initial conditions for this simulation (in terms of middle and deep soil layer resources and
 1003 biomass distribution) are taken from the results obtained at year 1915 of the previous simulation
 1004 using the reconstructed rainfall series in order to avoid, or at least minimise the effects of
 1005 adjusting to initial resource redistribution. Yao *et al.* (2006) reported a low (conservative)
 1006 grazing level applied on their quadrats over the period from 1915 to 2001 which was calculated
 1007 from averaged monthly stocking data that also accounted for changes in fence positions. We
 1008 therefore apply the same conservative grazing level following the management strategies
 1009 evolved by the Research Station over this period (Havstad *et al.*, 2006).

1010

1011 The centre-line averaged results are presented in figure 15, where it can be seen that the grass
 1012 biomass follows (but lags behind) the rainfall pattern (figure 5a). The average shrub biomass is
 1013 low, but remains above zero due to the grazing disturbance. The wetter years following the
 1014 1950s drought allow an accumulation of middle and deep-layer soil resources, and a trend for

1015 average connectivity values to increase is also evident. An increase in connectivity may indicate
 1016 that the spatial distribution tends towards a more open plant community (Okin *et al.*, 2001), or
 1017 that there is an increase in spatial organisation of the plant community.

1018
 1019 The spatial distributions of the grass and shrubs are plotted in figure 16. The increasingly
 1020 fragmented pattern in the grass distribution becomes apparent during the dry years of the 1920s
 1021 and 1930s, as a result of the persistence of shrubs (compare maps for 1925 and 1935). The
 1022 recovery process of both species is particularly well illustrated in the results for 1965, 1975 and
 1023 1985. During the drought of the 1950s, the shrubs persist as isolated spots, whereas the grass
 1024 survives as short horizontal bands (orientated across the slope). As the grass recovers, these
 1025 bands extend laterally, and then coalesce to form longer and more continuous bands, while the
 1026 shrubs recover to form isolated communities that are orientated in the direction of the driving
 1027 flux (shown also in figure 14). This pattern is caused by the shrubs' requiring longer connected
 1028 pathways than the grass (the length of this pathway is again a function biomass and resource
 1029 input) and because R&P diffusion around the shrub is vertical (through the soil layers) as well as
 1030 lateral.

1031
 1032 The contraction and recovery of biomass is also shown by plots of biomass and resource
 1033 concentration (figure 17) that are plotted for the centre-line. In 1925, cells with high grass
 1034 biomass are interspersed with cells with low biomass. Shrubs exist on the edges of these grass
 1035 peaks, and water resources are concentrated onto the grass patches. Following the drought and
 1036 subsequent recovery during the 1930s, the number of grass cells is reduced, and in some places,
 1037 the shrubs have recovered to higher biomass levels than those observed for the grass. This

1038 recovery happens where the vectors move excess resource to a patch, which then loses some
 1039 biomass by drought or disturbance.

1040
 1041 During the severe drought of the 1950s, most of the biomass is lost except in the locations that
 1042 contained the highest biomass prior to the drought, and recovery initiates from these cells during
 1043 the 1960s. From this point on, the surviving biomass exists at higher concentrations, but in
 1044 fewer cells. Resources are concentrated in patches, which are interesting in a number of ways.
 1045 First, the concentrations are at levels that greatly exceed the resource input to the grid. Secondly,
 1046 they exceed the resource requirement of the grass. Thirdly, they also exceed the level that can be
 1047 consumed, even under the maximum growth rate for the grass (table 4). The concentration of
 1048 resources into patches supports a higher biomass than the same resource could support, were the
 1049 grass more homogenously distributed (as suggested by Aguiar and Sala, 1999; Barbier *et al.*,
 1050 2006). It can be noted that the biomass on the up-gradient edge of the band is higher than at any
 1051 point earlier in the simulation on a more uniform grid (as observed by Ludwig *et al.*, 2005).

1052
 1053 Inspection of the resource transects show that for similar (and fairly static) banding patterns, the
 1054 distribution of subsurface resources can be very different from each other. On the transect
 1055 (figure 18), the shrubs are ultimately suppressed by the grass, but elsewhere in the grid, some
 1056 shrubs cells are able to survive. Two conditions are met by the surviving shrub cells. First, they
 1057 are those that contained a high pre-drought biomass. Secondly, these shrub cells are located on
 1058 the edge of a surviving grass band (McAuliffe, 1988). This pattern occurs because the grass
 1059 bands are able to concentrate resources to an extent that exceeds their maintenance and
 1060 maximum growth-rate requirements. The shrubs can capitalise on this excess resource, but

1061 elsewhere, any surviving shrubs that do not lie on the resource bands die off, even where the
 1062 initial pre-drought shrub density was high (Schlesinger *et al.*, 1990). Whether or not shrubs will
 1063 survive after the drought (during times when resource inputs are increased and grass populations
 1064 recover) is a little more complicated. Even with a sufficiently long connected path, shrub cells
 1065 will not persist in locations where resource movement is controlled by the grass, i.e.
 1066 predominantly lateral. In order to persist the shrubs also require sufficient vertical input of
 1067 resources that can be channelled to the deeper soil layers (a process documented in the literature,
 1068 e.g. Martinez_Meza and Whitford, 1996; and included in the model). This result points towards
 1069 the importance of different pathways of resources movement occurring within the shrub
 1070 populations – specifically a vertical connection between water input and deep layer soil
 1071 resources (which could be considered as a type of connectivity).

1072

1073 In figure 18, the experimental data of Yao *et al.* (2006) are compared to simulated data from
 1074 cells selected manually from the transect of figure 17 that show a similar response to Yao *et al.*'s
 1075 data. The selected simulated responses show markedly similar trends, although the specific
 1076 values are different. This difference most probably arises because the experimental data
 1077 measured the area covered by a plant in each quadrat to evaluate basal cover, whereas the model
 1078 computes biomass.

1079 In the first example (figure 18a), where grass is lost during the early 1920s, the biomass of this
 1080 cell was lower than in neighbouring cells, and so biomass was shed from this cell very quickly.
 1081 A number of upslope cells subsequently became connected, and the biomass increased slightly as
 1082 a result of the extra resource input, before becoming extinct in 1924. After grass was lost from
 1083 this cell, the number of upslope connected cells continued to increase. Without any nearby

1084 surviving patches to exchange resource and propagules with, this cell remained empty for the
 1085 remainder of the simulation (figure 17).

1086
 1087 The second example (figure 18b) shows the biomass response of a cell where grass was lost
 1088 during the drought of the 1950s. Initially, this cell was towards the downslope edge of a grass
 1089 patch. Grass persisted for a time on this patch while upslope cells were connected to the vectors.
 1090 A consequence of this increased connectivity was that new growth appears to have occurred on
 1091 the upslope edge of this cell and biomass was progressively lost on the downslope edge of this
 1092 band during periods of climatic stress. During the drought of the 1950s, there was insufficient
 1093 resource added by precipitation to sustain the biomass in this cell and the additional resource that
 1094 was delivered to the patch by the vectors was consumed by upslope biomass before it could
 1095 reach the cell. Consequently, grass in this location was lost. It is worth noting that towards the
 1096 end of the simulation, the grass in this cell is able to recover most likely because this simulation
 1097 does not account for concomitant soil degradation, which almost certainly occurs. This recovery
 1098 is not observed in the data of Yao *et al.* (2006).

1099
 1100 The third example (figure 18c) shows a cell where grass was lost immediately following the
 1101 1950s drought. This cell is located immediately upslope of the cell in the second example, and
 1102 the processes applying in the two locations are the same. The difference is that this cell is
 1103 located further towards the centre of the grass patch and survives for longer. After the drought,
 1104 the upslope biomass was able to increase quickly in response to higher rainfall inputs and higher
 1105 water input by the vectors. This rapid growth inhibits the recovery of this cell and ultimately

1106 grass is lost in this location, immediately following the drought. Elsewhere in the grid, an
 1107 upslope shrub cell has the same effect, in that it interrupts the connectivity to the grass cell.

1108
 1109 The final example (figure 18d) shows a cell where grass has survived throughout the duration of
 1110 the simulation. This cell exists near the upslope edge of a grass patch (*cf.* Montana, 1992) and as
 1111 such it has received high water inputs from upslope connected cells, even during the droughts,
 1112 and its connectivity was not interrupted following the drought.

1113
 1114 Yao *et al.* (2006) reported data from 98 1-m² quadrats in the Jornada, and reported that black
 1115 grama became locally extinct on 21% of these plots prior to the 1950s drought, 39% during the
 1116 drought, and 30% after the drought, and the grass persisted throughout the time of study in the
 1117 remaining 10% of quadrats that were examined. Simulation d generated biomass data in 1-m²
 1118 cells, using a parameterization consistent with the location of the data of Yao *et al.*. In
 1119 simulation d, black grama became locally extinct in 26% of the cells prior to the 1950s drought,
 1120 38% during the drought, 20% after the drought, and black grama persisted in 16% of the cells.
 1121 The simulated and measured results show remarkably good agreement with each other, with the
 1122 largest differences occurring after the 1950s drought where simulation d overpredicts the number
 1123 of cells where grass survives. This overprediction may be because there is no mechanism within
 1124 the model that would allow for a spatial location to be degraded by hillslope processes during a
 1125 disturbance and therefore inhibit biomass recovery (Montana, 1992; Abrahams *et al.*, 1995).

1126

1127

1128 **7. Discussion**

1129 In the early part of this paper, we have argued that, though modelling can provide unique insights
 1130 into understanding the dynamics of the patchiness of desert vegetation, such insights are valuable
 1131 only if models yield testable predictions and if the models are firmly grounded in, and
 1132 compatible with, empirical data. That is not to say that empirical data are uncontestable. All
 1133 data are collected within a conceptual framework, and it may be that modelling will yield results
 1134 that lead us to challenge that conceptual framework. Even so, such a challenge is only valid
 1135 where the model makes explicit reference to that conceptual framework. Against that argument,
 1136 we have developed a numerical model for the dynamics of desert ecosystems within the
 1137 conceptual framework of connectivity, and we have parameterized the model for implementation
 1138 in a specific desert setting where the available data set for parameterization is particularly rich.
 1139 In this section, we discuss the model output in terms of its ability to generate testable hypotheses.
 1140
 1141 Four testable hypotheses emerge from our modelling of shrub invasion of grasslands. First, our
 1142 results show that, contrary to Thornes and Brandt (1993), rainfall variability does not enhance
 1143 shrub invasion. Instead, because the annual re-growth rate of grass rate is higher than that of
 1144 shrubs it recovers faster and is thus able to suppress shrub invasion by re-establishing its control
 1145 on resource redistribution. This result suggests the hypothesis that the propensity of a grassland
 1146 for shrub invasion is a function of the relative growth rate of the two. It should be noted,
 1147 however, as a caveat to this hypothesis there may be extremes of drought beyond those tested
 1148 here under which such a control breaks down.

1149

1150 Secondly, Yao *et al* (2006) hypothesized that control by transport processes on local resource
 1151 and propagule distribution may somehow explain the spatial variation in grass survival within
 1152 the Jornada Basin. Our analysis suggests that the timing of grass loss depends on both initial cell
 1153 biomass and connectivity properties. Low biomass density patches will always become extinct
 1154 first during times of resource shortage. During prolonged periods of drought, biomass on the
 1155 downslope edge of a patch is the most vulnerable, particularly if a shrub exists nearby, because
 1156 the shrub will channel some of the resource input to deeper layers and so a longer connected
 1157 pathway or a greater resource input is required to supply the grass cell than would be required if
 1158 the shrub were not located nearby. Cells that lose biomass tend to have low numbers of upslope
 1159 connected cells. Patches that survive drought conditions are those that are located near the
 1160 upslope edge of the patch, where there is a high degree of connectivity in up-vector cells. These
 1161 results lead us to suggest the hypothesis that changes in the values of connectivity for grassland
 1162 indicate conditions where it would be particularly vulnerable to a disturbance externality.

1163
 1164 Thirdly, our results lead us to hypothesize that when the grass species becomes established and
 1165 forms a stable community, it is able to control the resource and propagule movement within the
 1166 landscape to suit its own survival strategy. In the case of pure stands of grass, this hypothesis
 1167 means that the resource distributions coincide with the scale of grass plants, and underlying
 1168 resources in deeper layers are at a minimum (Müller *et al.*, 2008). When conditions change to
 1169 allow deeper layer resources to accumulate, shrub invasion into a grass stand can occur. This
 1170 hypothesis therefore predicts a vertical resource gradient should occur beneath shrub
 1171 communities, whereas the resource gradient surrounding grass patches should be predominantly
 1172 lateral. These predictions are supported at least qualitatively by the results of Schlesinger *et al.*

1173 (1996). The simulations also suggest that a vertical resource profile could be an independent
 1174 means by which the islands of fertility model can be tested. Furthermore, it has been argued that
 1175 under conditions (that are usually driven by a strong externality) where the dominant grass
 1176 species has collapsed, permanent changes in soil condition (such as caused by erosion, as noted
 1177 by Westoby *et al.*, 1989; Abrahams *et al.*, 1995; Li *et al.*, 2007) would become much more
 1178 important and are likely to inhibit future vegetation establishment in eroded areas (Wainwright *et*
 1179 *al.*, 2000; Okin *et al.*, 2006; Li *et al.*, 2009). Mauchamp *et al.*, (1993) theorised that stripes are
 1180 controlled by different recruitment histories, and that on a landscape scale it is the successive
 1181 die-back and regrowth that controls stripes. Because our model only allows these different
 1182 recruitment processes to occur in response to resource inputs, and because we are, nevertheless,
 1183 able to generate plausible results, it leads to the hypothesis that resources rather than changes in
 1184 soil conditions *per se* that are the primary controlling factors.

1185
 1186 Fourthly, the distinct differences in our modelling results between diffusion-advection and
 1187 advection-only simulations (see Appendix) lead us to hypothesize that the balance of these two
 1188 sets of processes and the nature of diffusion play a large part in controlling vegetation behaviour.
 1189 It would be possible to test this hypothesis by conducting field experiments in which the ability
 1190 of vectors to move resources and propagules in these ways was examined.

1191
 1192 That our model is able to generate specific testable hypotheses is due to the specific
 1193 parameterization that we have been able to provide for the data-rich Jornada Basin. However,
 1194 the relationship between the utility of a model and the available data for parameterization is not a
 1195 simple one. Where such data do not exist, a model may be used to identify specific data needs

1196 for testing of hypotheses. Sensitivity of model output to particular parameters can drive
 1197 empirical research just as much as the results of empirical research can lead a model to produce
 1198 testable hypotheses. For example, key parameters in understanding the process of shrub invasion
 1199 and the potential for its reversal are the rates of establishment and mortality for creosotebush.
 1200 Current estimates for these parameters are derived from other modelling studies (Peters, 2002a),
 1201 suggesting that further empirical work is needed to constrain the potential values of parameters
 1202 to which the model is highly sensitive.

1203

1204

1205 **8. Conclusions**

1206 In this paper, a modelling framework that explicitly considers spatial interactions among
 1207 multiple vegetation types and multiple resources has been applied to the analysis of ecosystem
 1208 change in deserts. The model is designed to explore causes of spatial complexity as well as
 1209 predict specific responses to a variety of endogenous and exogenous disturbances. This
 1210 contribution differs from previous work in that it rests on a sound process-based understanding
 1211 and data that has both a clear physical meaning and can be measured in the field. Both abiotic
 1212 and biotic processes have been considered in greater detail than previous modelling studies,
 1213 while maintaining a level of parsimony that means that parameter uncertainty is unlikely to
 1214 drown out the effects of the processes under investigation.

1215

1216 A general modelling framework has been developed, and specific implementation of this model
 1217 was employed to evaluate the framework against data that has been obtained from field studies.
 1218 In doing so, it is noted that even with the simplifications made, the model was able to closely

1219 match measured conditions at the field site, in terms of species response and the generation of
 1220 plausible patterns of vegetation loss. On this basis, the general framework can be considered to
 1221 have captured the key processes within the ecosystem and may make a useful contribution
 1222 towards understanding desert vegetation more straightforwardly.

1223
 1224 Rather than developing predictions of vegetation change under hypothetical future scenarios,
 1225 historical data have been used to retrodict grassland responses to climatic conditions. In doing
 1226 so, it was possible to compare the model results to current conditions, which comparison
 1227 provides a robust test of both the model and our understanding of how desert ecosystems operate.
 1228 Moreover, the approach has led to the generation of a number of testable predictions that can be
 1229 compared to other field data.

1230
 1231 The results suggest that the desert grasslands have been stable under historic conditions for three
 1232 reasons. First, the structure of the rainfall itself inhibits shrub invasion; secondly, the faster
 1233 growth rate tends to allow grass to outcompete with shrubs for available resources; and thirdly,
 1234 the banding patterns themselves are much more stable structures in semi-arid ecosystems than a
 1235 homogenous distribution of grass. During droughts, the resource that is input to connected cells
 1236 will flow onto the bands where it supports the patch biomass, and the length of the connected
 1237 cells and the length of grass bands are related. The resource distribution across patches remains
 1238 predominantly lateral, with little (or no) resource accumulated beneath the bands. The shrubs (in
 1239 this simulation) are only able to invade grass stands when a disturbance causes a grass plant to be
 1240 removed from a location where resource has accumulated, but the persistence of the shrub is also
 1241 a function of two types of connectivity. First, the length of the connected pathways to the shrub

1242 must be longer than the connected pathways to the grass plant, and secondly, shrubs will only
 1243 survive in locations where they are able to develop a pronounced vertical distribution of
 1244 resources beneath them. This difference would suggest that lateral accumulations of resource
 1245 around a shrub indicate that recent climate conditions have caused a great degree of R&P
 1246 movement along newly emerged connected pathways, but the absence of that accumulation
 1247 points towards vector operation along more stable connected pathways. The results point to the
 1248 introduction of cattle grazing, and specifically overgrazing, as the cause of the historical shrub
 1249 invasion.

1250

1251 The agreement of the model results with experimental studies indicates that this method has
 1252 merit and is worth pursuing further. It is acknowledged that the implementation presented here
 1253 far from perfect in two significant respects. First, we have, for instance, used a number of linear
 1254 relationships which would not be appropriate in a more general implementation, and many of the
 1255 surrogate data used here (with respect to the relationships between vectors and resource
 1256 movement) ought to be parameterised more fully. In particular, experimental data aiming to
 1257 quantify the redistribution of laterally-transported resources at the end of a connected pathway
 1258 (i.e. Equation (6)) is deserving of attention. The lack of detail supported by field research in
 1259 these factors means that while we can see that the connected pathways must be longer to enable
 1260 shrubs to survive than for grass plants, we are not able to quantify them with any confidence.
 1261 Secondly, some of our model results are unrealistic. For example, the accumulation of nitrogen
 1262 in the mid and deep soil layers is unrealistic. As was pointed out in the model parameterization
 1263 (5.1.4), there is a dearth of suitable data for this parameterization. Our results suggest that this
 1264 dearth of data is a significant limitation on our current understanding. The application of the

1265 model to a very specific implementation was worthwhile in order to establish that the general
 1266 framework produces plausible results, and to inform future experimental work that may obtain
 1267 data in the form required to establish the causal factors that lead to ecosystem changes. On the
 1268 basis of this work, vertical and lateral connectivity are key emergent properties of the system
 1269 which both control its behaviour and provide indicators of its state. If these predictions are
 1270 shown to be compatible with actual conditions, the model presented here will provide a more
 1271 certain approach towards preventing further semi-arid grassland degradation.

1272

1273

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1278 **REFERENCES**

1279 Abrahams, A. D., Parsons, A. J., 1991. Relation between infiltration and stone cover on a
 1280 semiarid hillslope, southern Arizona., J. Hydrol., 122, 49–59.

1281 Abrahams, A.D., Parsons, A.J., Wainwright, J., 1995. Effects of vegetation change on interrill
 1282 runoff and erosion, Walnut Gulch, southern Arizona. Geomorph. 13(1-4), 37-48.

1283 Abrahams, A.D., Parsons, A.J., Wainwright, J., 2003. Disposition of rainwater under
 1284 creosotebush. Hydrol. Proc. 17, 2555-2566.

1285 Ackerman, T.L., Bamberg, S.A., 1974. Phenological studies in the Mojave Desert at Rock
 1286 Valley. In: H. Lieth (Editor), Ecological studies; Analysis and synthesis Springer-Verlag,
 1287 New York, pp. 215-226.

- 1288 Aguiar, M.R., Sala, O.E., 1999. Patch structure, dynamics and implications for the functioning of
 1289 arid ecosystems. *Trends in Ecology and Evolution*. 14, 273-277.
- 1290 Allred, K.W., 1996. New Mexico's Natural Heritage: Biological diversity in the land of
 1291 enchantment. *New Mexico Acad. Sci. Albuquerque*, pp 168-231
- 1292 Archer, S., Schimel, D.S., Holland, E.A., 1995. Mechanisms of shrubland expansion: land use,
 1293 climate or CO₂. *Climatic Change*. 29, 91-99.
- 1294 Baez, S., Fargione, J., Moore, D.I., Collins, S.L., Gosz, J.R., 2007. Atmospheric nitrogen
 1295 deposition in the northern Chihuahuan desert: Temporal trends and potential consequences.
 1296 *J. Arid Environ.* 68, 640-651.
- 1297 Bainbridge, D.A., Virginia, R.A., 1990. Restoration in the Sonoran Desert of California.
 1298 *Restoration and Management Notes*, 8, 3-14.
- 1299 Barbier, N., Couteron, P., Lejoly, J., Deblauwe, V., Lejeune, O., 2006. Self-organized vegetation
 1300 patterning as a fingerprint of climate and human impact on semi-arid ecosystems. *J.*
 1301 *Ecology*. 94, 537-547.
- 1302 Barbour, M.G., 1969. Age and space distribution of the desert shrub *Larrea divaricata*. *Ecology*.
 1303 50, 679-685.
- 1304 Bartley, R., Roth, C.H., Ludwig, J.A., McJannet, D., Leidloff, A.C.,
 1305 Corfield, J., Hawdon, A., Abbot, B., 2006. Runoff and erosion from
 1306 Australia's tropical semi-arid rangelands: Influences of ground cover for
 1307 differing space and time scales. *Hydrological Processes* 20: 3317-3333
- 1308 Bracken LJ, Croke J. 2007. The concept of hydrological connectivity and its contribution to
 1309 understanding runoff-dominated geomorphic systems. *Hydrological Processes* 21, 1749–
 1310 1763

- 1311 Brazier, R.E., Parsons, A.J., Wainwright, J., Powell, D.M., Schlesinger, W.H., 2007. Upscaling
 1312 understanding of nitrogen dynamics associated with overland flow in a semi-arid
 1313 environment. *Biogeochemistry*, 82, 265-278.
- 1314 Brisson, J. and Reynolds, J.F. 1994 'The effect of neighbors on root distribution in a
 1315 creosotebush (*Larrea tridentata*) population', *Ecology* **75**, 1693–1702.
- 1316 Bromley, J., Brouwer, J., Barker, A.P., Gaze, S.R., Valentin, C., 1997. The role of surface water
 1317 distribution in an area of patterned vegetation in a semi-arid environment, south west
 1318 Niger. *J. Hyd.* 198, 1-29.
- 1319 Brown, D.E., 1982. Semi-desert grassland. In: Brown, David E., ed. *Biotic communities of the*
 1320 *American Southwest--United States and Mexico. Desert Plants.* 4(1-4), 123-131.
- 1321 Brown, D.W., Gersmehl, P.J., 1985. Migration models for grasses in the American midcontinent.
 1322 *Annals Association Am. Geographers*, 75, 383-394.
- 1323 Buffington, L.C., Herbel, C.H., 1965. Vegetational changes on a semi-desert grassland range
 1324 from 1858 to 1963. *Ecol. Monogr.* 35, 139-164.
- 1325 Bugmann, H.K.M., Solomon, A.M., 1995. The use of a European forest model in North
 1326 America: a study of ecosystem response to climate gradients. *J. Biogeog.* 22, 477-484.
- 1327 Burgess, T.L., 1995. Desert grassland, mixed shrub savanna, shrub steppe or semidesert scrub?
 1328 The dilemma of coexisting forms. In: M.P. McClaran and T.R. Van Devender (eds.), *The*
 1329 *desert grassland.* The University of Arizona Press, Tuscon, pp. 31-67.
- 1330 Burke, I.C., Schimwl, D.S., Yonker, C.M., Parton, W.J., Joyce, L.A., and Lauenroth, W.K.,
 1331 1990. Regional modelling of grassland biogeochemistry using GIS. *Landscape Ecology.* 4,
 1332 45-54.

- 1333 Campbell, R.S., 1931. Plant succession and grazing capacity on clay soils in southern New
 1334 Mexico. *J. Agri. Res.* 43(12), 1027-1051.
- 1335 Canfield, R.H., 1948. Perennial grass composition as an indicator of condition of Southwestern
 1336 mixed grass ranges. *Ecology*, 29, 190-204.
- 1337 Casenave, A., Valentin, C., 1992. A runoff capability classification-system based on surface-
 1338 features criteria in semiarid areas of West Africa. *J. Hydrol.* 130, 231–249.
- 1339 Casper, B.B., Jackson, R.B., 1997. Plant competition below ground. *Ann. Rev. Ecological*
 1340 *Systems* 28, 545-570.
- 1341 Castellanos, A.E., Molina, F.E., 1990. Differential survivorship and establishment in
 1342 *Simmondsia chinensis* (jojoba). *J. Arid Environ.* 19, 65-76.
- 1343 Chandrasekhar, S. (Editor) 1961. *Hydrodynamic and Hydromagnetic Stability*. Springer-Verlag,
 1344 Berlin.
- 1345 Charley, J.L., West, N.E., 1975. Plant-induced soil chemical patterns in some shrub-dominated
 1346 semi-desert ecosystems of Utah. *J. Ecology*.
- 1347 Chew, R., Chew, A.E., 1970. Energy relationships of the mammals of a desert shrub (*Larrea*
 1348 *tridentata*) community. *Ecol. Monogr.* 40, 1-21.
- 1349 Clark, C.M., Tilman, D., 2008. Loss of plant species after chronic low-level nitrogen deposition
 1350 to prairie grasslands. *Nature*. 451, 712-715.
- 1351 Clos-Arceuduc, M., 1956. Etude sur photographies aeriennes d'une formation vegetale
 1352 sahelienne: la brousse tigre. *Buletin de L'IFAN Serie A* 18, 678-684.
- 1353 Cody, M.L., 1986. Spacing patterns in the Mojave Desert plant communities: near neighbour
 1354 analysis. *J. Arid Environ.* 11, 199-217.

- 1355 Coffin, D.P., Lauenroth, W.K., 1990. A gap dynamics simulation model of succession in the
 1356 shortgrass steppe. *Ecol. Model.* 49, 229-266.
- 1357 Comrie, A.C., Broyles, B. 2002. Variability and spatial modelling of fine-scale precipitation data
 1358 for the Sonoran Desert of south-west Arizona. *J. Arid Environ* 50, 573-592
- 1359 Cornet, A., Delahoume, J.P., Montana, C. (Editors), 1988. Dynamics of striped vegetation
 1360 patterns and water balance in the Chihuahuan Desert. SBP Academic Publishing, The
 1361 Hague, pp 221-231
- 1362 Couteron, P., Lejeune, O., 2001. Periodic spotted patterns in semi-arid vegetation explained by a
 1363 propagation-inhibition model. *J. Ecology.* 89, 616-628.
- 1364 Dakos, V., Kéfi, S., Rietkerk, M., van Nes, E.M., Scheffer, M. 2011. Slowing down in spatially
 1365 patterned ecosystems at the brink of collapse. *The American Naturalist*, 177, E153-E166.
- 1366 d'Arigo, R.D., Jacoby, G.C., 1992. A tree-ring reconstruction of New Mexico winter
 1367 precipitation and its relation to El Nino/Southern Oscillation events. In: H.F. Diaz and V.
 1368 Markgraf (eds.), *El Nino. historical and Paleoclimatic Aspects of the Southern Oscillation.*
 1369 Cambridge University Press, Cambridge, pp. 7-28.
- 1370 Deblauwe, V., Barbier N., Couteron P., Lejeune O., Bogaert J., 2008. The global biogeography
 1371 of semi-arid periodic vegetation patterns. *Global Ecology and Biogeography* 17, 715–723.
- 1372 d'Herbes, J.M., Valentin, C., Tongway, D.J. (eds.), 2001. Banded vegetation patterning in arid
 1373 and semiarid environments; ecological processes and consequences for management.
 1374 Springer, Hiedelberg, 1-19 pp.
- 1375 D'Odorico, P., Laio, F., Ridolfi, L., 2006. Patterns as indicators of productivity enhancement by
 1376 facilitation and competition in dryland vegetation. *Journal of Geophysical Research* 111, G03010.

- 1377 Dunkerley, D.L., Brown, K.J., 1999. Banded vegetation near Broken Hill, Australia: significance
 1378 of surface roughness and soil physical properties. *Catena*. 37(1-2), 75-88.
- 1379 Furukawa, K., Imai, K., Kurashige, M., 2000. Simulated effect of box size and wall on porosity
 1380 of random packings of spherical particles. *Acta Mechanica* .140, 219-231.
- 1381 Gallardo, A., Schlesinger, W.H., 1992. Carbon and nitrogen limitations of soil microbial biomass
 1382 in desert ecosystems. *Biogeochemistry*, 18, 1-17.
- 1383 Gardener, J.L., 1950. The effects of thirty years of protection from grazing in desert grasslands
 1384 *Ecology*, 31, 44-50.
- 1385 Gibbens, R.P., McNeely, R.P., Havstad, K.M., Beck, R.F., Nolen, B., 2005. Vegetation changes
 1386 in the Jornada Basin from 1858 to 1998. *J. Arid. Environ.* 61, 651-668.
- 1387 Gibbens, R.P., Lenz, J.M., 2001. Root system of some Chihuahuan Desert plants. *J. Arid*
 1388 *Environ.* 49, 221-263.
- 1389 Gillett, J., 1941. The plant formations of western British Somaliland and the Harar province of
 1390 Abyssinia. *Kew Bull.* 2, 37-75.
- 1391 Goldberg, D.E., Turner, R.M., 1986. Vegetation change and plant demography in permanent
 1392 plots in the Sonoran desert. *Ecology*. 67, 695-712.
- 1393 Gosz, R.J., Gosz, J.R., 1996. Species interactions on the biome transition zone in New Mexico:
 1394 response of blue grama (*Bouteloua gracilis*) and black grama (*Bouteloua eripoda*) to fire
 1395 and herbivory. *J. Arid Environ.* 34, 101-114.
- 1396 Greene, R. S. B., Kinnell P. I. A., Wood J. T., (1994) Role of plant cover and stock trampling on
 1397 runoff and soil erosion from semi-arid wooded rangelands. *Aust. J. Soil. Res.* 32, 954-973.
- 1398 Greig-Smith, P., 1979. Pattern in vegetation. *J. Ecology*. 67, 755-779.

- 1399 Grimm, V., Railsback, S.F., 2005. Individual-Based Modelling and Ecology. Princeton
 1400 University Press, Princeton.
- 1401 Havstad, K.M., Huenneke, L.F., Schlesinger, W.H. (eds.), 2006. Structure and Function of a
 1402 Chihuahuan Desert Ecosystem: The Jornada Basin Long-Term Ecological Research Site.
 1403 Oxford University Press, New York.
- 1404 Herbel, C.H., Ares, F.N., Wright, R.A., 1972. Drought effects on a semidesert grassland range.
 1405 Ecology, 53, 1084-1093.
- 1406 Higgins, S.I., Richardson, D.M., Cowling, R.M., 1996. Modelling invasive plant spread: the role
 1407 of plant-environment interactions and model structure. Ecology. 77, 2043-2054.
- 1408 Hillel, D., 2004. Introduction to Soil Physics. Elsevier.
- 1409 HilleRisLambers, R., Rietkerk, M., van den Bosch, F., Prins, H.H.T., de Croon, H., 2001.
 1410 Vegetation Pattern formation in semi-arid grazing systems. Ecology. 82, 50-61.
- 1411 Hooper, D.U., Johnson, L., 1999. Nitrogen limitation in dryland ecosystems: Responses to
 1412 geographical and temporal variation in precipitation. Biogeochemistry, 46, 247-293.
- 1413 Humphrey, R.R., 1958. The desert grassland: a history of vegetational change and an analysis of
 1414 causes. Botanical Rev. 24, 164-193.
- 1415 Humphrey, R.R., Mehrhoff, L.A., 1958. Vegetation changes on a southern Arizona grassland
 1416 range. Ecology. 39, 720-726.
- 1417 Istanbuluoglu, E. and Bras, R.L., 2006. On the dynamics of soil moisture, vegetation, and
 1418 erosion: Implications of climate variability and change. Water Resour. Res. 42.
- 1419 Jornada LTER 2010 Jornada Basin Long Term Ecological Research. <http://jornada->
 1420 www.nmsu.edu/ [last accessed 21st July 2010]

- 1421 Kéfi, S., Rietkerk, M., van Baalen, M., and Loreau, M. 2007. Local facilitation, bistability and
 1422 transitions in arid ecosystems. *Theoretical Population Biology* 71:367–379.
- 1423 Klausmeier, C.A., 1999. Regular and Irregular Patterns in Semiarid Vegetation. *Science* 284,
 1424 1826-1828.
- 1425 Kot, M., 2001. *Elements of Mathematical Ecology*. Cambridge University Press.
- 1426 Lange, R.T., Nicholson, A.D., Nicholson, D.A., 1984. Vegetation management of chenopod
 1427 rangelands in South Australia. *Aust. Rangeland J.* 6, 46-54.
- 1428 Leach, A.R., 2001. *Molecular Modelling: Principals and Applications*. Prentice Hall.
- 1429 Lefever R., Barbier, N., Couteron, P., and Lejeune, O., 2009. Deeply gapped vegetation patterns:
 1430 on crown/root allometry, criticality and desertification. *Journal of Theoretical Biology*
 1431 261, 194–209.
- 1432 Loreau, M., Naeem, S., Inchausti, P., 2002. *Biodiversity and ecosystem functioning: synthesis*
 1433 *and perspectives* Oxford University Press.
- 1434 Li, J., Okin, G. S., Hartman, L. J., Epstein, H. E., 2007. Quantitative assessment of wind erosion
 1435 and soil nutrient loss in desert grasslands of southern New Mexico, USA.
 1436 *Biogeochemistry* 85, 317-332.
- 1437 Ludwig, J.A., Wilcox, B.P., Breshears, D.D., Tongway, D.J., Imeson, A.C., 2005. Vegetation
 1438 patches and Runoff-Erosion as Interacting Ecohydrological Processes in Semiarid
 1439 Landscapes. *Ecology*. 86, 288-297.
- 1440 MacFadyan, W.A., 1950. Vegetation patterns in the semi-desert plains of British Somaliland.
 1441 *Geographical Journal* 116, 199–210.

- 1442 Maneta, M.P., Schnabel, S., Wallender, W.W., Panday, S., Jetten, V., 2008. Calibration of an
 1443 evapotranspiration model to simulate soil water dynamics in a semiarid rangeland.
 1444 Hydrological Proc., 22.
- 1445 Michaelides, K., Lister, D., Wainwright, J., Parsons, A.J., 2012. Linking runoff and erosion
 1446 dynamics to nutrient fluxes in a degrading dryland landscape. Journal of Geophysical
 1447 Research – Biogeosciences 117, G00N15, doi:10.1029/2012JG002071.
- 1448 Martinez-Meza, E., Whitford, W.G., 1996. Stemflow, throughfall and channelization of
 1449 stemflow by roots in three Chihuahuan desert shrubs. J. Arid Environ. 32, 271-287.
- 1450 Mauchamp, A., Montana, C., Lepart, J., Rambal, S., 1993. Ecotone dependent recruitment of a
 1451 desert shrub, *Flourensia cernua*, in vegetation stripes. Oikos. 68, 107-116.
- 1452 May, R.M., 1977. Thresholds and breakpoints in ecosystems with a multiplicity of stable states.
 1453 Nature. 269, 477.
- 1454 McAuliffe, J.R., 1988. Markovian dynamics of simple and complex desert plant communities.
 1455 The Am. Naturalist, 131, 459-490.
- 1456 McClaran, M. P., Van Devender, T. R., 1995. The Desert Grassland. Tuscon University of
 1457 Arizona Press
- 1458 McPherson, G.R., Wright, H.A., Wester, D.B., 1988. Patterns of shrub invasion in semi-arid
 1459 Texas grasslands. Am. Midland Naturalist. 120, 391-397.
- 1460 Meinhardt, H., 1982. Models of Biological Patterns Formation. Academic, New York.
- 1461 Miller, R.F., Donart, G.B., 1979. Response of *Bouteloua eriopoda* (Torr.) Torr. and *Sporobolus*
 1462 *flexuosus* (Thurb.) Rhbd to season of defoliation. J. Range Manage. 32, 449-452.
- 1463 Monger, H.C., Bestelmeyer, B.T. 2006. The soil-geomorphic template and biotic change in arid
 1464 and semi-arid ecosystems. J. Arid Environ. 65, 207-218.

- 1465 Montana, C., 1992. The colonization of bare areas in two phase mosaics of an arid ecosystem. *J.*
 1466 *Ecology*. 80, 315-327.
- 1467 Morgan R.P.C (1996) *Soil erosion and conservation*, second edition, Harlow: Longman.
- 1468 Müller, E.N., Wainwright, J., Parsons, A.J., 2007. The impact of connectivity on the modelling
 1469 of water fluxes in semi-arid shrubland environments. *Water Resources Research* 43: doi
 1470 10-1029/2006WR005006.
- 1471 Müller, E.N., Wainwright, J., Parsons, A.J., 2008. Spatial variability of soil and nutrient
 1472 parameters within grasslands and shrublands of a semi-arid environment, Jornada Basin,
 1473 New Mexico. *Ecohydrology*. 1, 3-12.
- 1474 Murray, J.D., 1989. *Mathematical Biology*. Springer-Verlag, Berlin.
- 1475 Nash, M.S., Jackson, E., Whitford, W.G., 2004. Effects of intense, short-duration grazing on
 1476 micro-topography in a Chihuahuan Desert grassland. *J. Arid Environ.* 56(3), 383-393.
- 1477 Neilson, R.P. (1986). High-resolution climatic analysis and Southwest biogeography. *Science*.
 1478 232, 27-34.
- 1479 Nelson, E.W., 1934. *The influence of precipitation and grazing upon black grama grass range*,
 1480 US Department of Agriculture, Washington, DC.
- 1481 Nicholson, S.E., 1979. Revised Rainfall series for the West African Subtropics. *Monthly*
 1482 *Weather Rev.* 107, 620-623.
- 1483 Nicholson, S.E., 1981. Rainfall and atmospheric circulation during drought periods and wetter
 1484 years in West Africa. *Monthly Weather Rev.* 109, 2191-2208.
- 1485 Noy-Meir, I., 1973. Desert Ecosystems: Environment and Producers. *Ann. Rev. Ecology*. 4, 25-
 1486 51

- 1487 Okin, G.S., Gillette, D.A., 2001. Distribution of vegetation in wind-dominated landscapes:
 1488 Implications for wind erosion modelling and landscape processes. *J. Geophys. Res.* 106,
 1489 9673-9683.
- 1490 Okin, G.S., Murray, B., Schlesinger, W.H., 2001. Degradation of sandy arid shrubland
 1491 environments: observations, process modelling and management implications. *J. Arid*
 1492 *Environ.* 47,123-144.
- 1493 Okin, G. S., Herrick, J. E., Gillette, D. A., 2006. Multiscale controls on and consequences of
 1494 aeolian processes in landscape change in arid and semiarid environments. *Journal of Arid*
 1495 *Environments* 65, 253-275.
- 1496 Okin, G.S., D'Odorico, P.D., Archer, S., 2009. Impact of feedbacks on Chihuahuan desert
 1497 grasslands: transience and metastability driven by grass recruitment. *Global Change*
 1498 *Biology*.
- 1499 Okin, G.S., Parsons, A.J., Wainwright, J., Herrick, J.E., Bestelmeyer, B.T., Peters, D.P.C.,
 1500 Fredrickson, E.L., 2009. Do changes in connectivity explain desertification?. *BioScience*
 1501 59, 237-244.
- 1502 Osborne, C.P. and Beerling, D.J., 2006. Nature's green revolution: the remarkable evolutionary
 1503 rise of C₄ plants. *Phil. Trans. Royal Soc. London B Bio. Sci.* 361, 173-194.
- 1504 Parshall, T., Foster, D.R., Faison, E., MacDonald, D. and Hansen, B.C.S., 2003. Long-term
 1505 history of vegetation and fire in pitch pine-oak forests on Cape Cod, Massachusetts.
 1506 *Ecology.* 84, 736-748.
- 1507 Parsons, A.J., Abrahams, A.D., Simanton, J.R., 1992. Microtopography and soil-surface
 1508 materials on semi-arid piedmont hillslopes, southern Arizona. *J. Arid Environ.* 22, 107-
 1509 115.

- 1510 Parsons, A.J., Abrahams, A.D., Wainwright, J., 1996. Responses of interrill runoff and erosion
 1511 rates to vegetation change in southern Arizona. *Geomorph.* 14, 311-317.
- 1512 Parsons, A.J., Wainwright, J., Abrahams, A.D., Simanton, J.R., 1997. Distributed dynamic
 1513 modelling of interrill overland flow. *Hydrol. Proc.* 11, 1833-1859.
- 1514 Parsons, A.J., Wainwright, J., Powell, D.M., Kaduk, J., Brazier, R., 2004. A conceptual model
 1515 for determining soil erosion by water. *Earth Surf. Process. Landforms.* 29, 1293-1302.
- 1516 Parsons, A.J., Brazier, R.E., Wainwright, J., Powell, D., 2006a. Scale relations in hillslope runoff
 1517 and erosion. *Earth Surf. Process. Landforms.* 31, 1384-1393.
- 1518 Parsons, A.J., Abrahams, A.D., Luk, S.H., 2006b. Size characteristics of sediment in interrill
 1519 overland flow on a semiarid hillslope, Southern Arizona. *Earth Surf. Process. Landforms.*
 1520 16, 143-152.
- 1521 Parton, W.J., 1978. Dynamics of C, N, P and S in grassland soils: a model. *Biogeochemistry* 5,
 1522 109-131.
- 1523 Paulsen, H.A. Jr., Ares, F.N., 1962. Grazing values and management of black grama and tobosa
 1524 grasslands and associated shrub ranges of the Southwest. *Tech. Bull. No. 1270.*
 1525 Washington, DC: U.S. Department of Agriculture, Forest Service, pp 56
- 1526 Peters, D.P.C., 2002a. Plant species dominance at a grassland-shrubland ecotone: an individual-
 1527 based gap dynamics model of herbaceous and woody species. *Ecol. Model.* 152, 5-32.
- 1528 Peters, D.P.C., 2002b. Recruitment potential of two perennial grasses with different growth
 1529 forms at a semiarid-arid transition zone. *Am. J. Botany.* 89(10), 1616-1623.
- 1530 Peters, D.P.C., Gosz, J.R., Pockman, W.T., Small, E.E., Parmenter, R., Collins, S.L., Polley, E.,
 1531 Johnson, H.B., Derner, J D., 2002. Soil- and plant-water dynamics in a C3/C4 grassland

- 1532 exposed to a subambient to superambient CO₂ gradient. *Global Change Biology*. 8, 1118–
 1533 1129.
- 1534 Peters, D.P.C., Bestelmeyer, B.T., Herrick, J.E., Fredrickson, E., Monger, H.C., Havstad, K.M.,
 1535 2006a. Disentangling complex landscapes: new insights to forecasting arid and semiarid
 1536 system dynamics. *Bioscience*. 56, 491-501.
- 1537 Peters, D.P.C., Gosz, J.R., Pockman, W.T., Small, E.E., Parmenter, R., Collins, S.L., Muldavin,
 1538 E., 2006b. Integrating patch and boundary dynamics to understand and predict biotic
 1539 transistions at multiple scales. *Landscape Ecology*. 21, 19-23.
- 1540 Pockman, W.T., Sperry, J.S., 1997. Freezing-induced xylem cavitation and the northern limit of
 1541 *Larrea tridentata*. *Oecologia*. 109,19-27.
- 1542 Rango, A., Chopping, M., Ritchie, J., Havstad, K., Kustas, W., Schmugge, T., 2000.
 1543 Morphological Characteristics of Shrub Coppice Dunes in Desert Grasslands of Southern
 1544 New Mexico derived from Scanning LIDAR Remote Sensing of Environ. 74, 26-44.
- 1545 Reynolds, J.F., Virginia, R.A., Kemp, P.R., de Soyza, A.G., Tremmel, D.C., 1999. Impact of
 1546 drought on desert shrubs: Effects of seasonality and degree of resource island
 1547 development. *Ecol. Monogr*. 69, 69-106.
- 1548 Rietkerk, M., Boerlijst, M.C., van Langevelde, F., HilleRisLambers, R., Van de Koppel, J.,
 1549 Kumar, L., de Roos, A.M., 2002. Self-organization of vegetation in arid ecosystems. *Am.*
 1550 *Naturalist*, 160, 524-530.
- 1551 Rietkerk, M.F., Van der Koppel, J., 1997. Alternate stable states and threshold effects in semi-
 1552 arid grazing systems. *Oikos* 79, 69-76.
- 1553 Romney, E.M., Wallace, A., Hunter, B., 1989. Pulse establishment of woody shrubs of denuded
 1554 Mojave Desert land. In: W. A, E.D. McArthur and M.R. Haferkamp (eds), *Symposium on*

- 1555 shrub ecophysiology and biotechnology. US Department of Agriculture, forest service,
 1556 Intermountain Research station, Logan, UT, pp. 54-57.
- 1557 Rovinsky, J.B., 1987. Twinkling Patterns and Diffusion-Induced Chaos in a Model of the
 1558 Belousov-Zhabotinsky Chemical Medium. *J. Phys. Chem.* 91.
- 1559 Seghier, J., Galle, S., Rajot, J.L., Ehrmann, M., 1997. Relationships between soil moisture and
 1560 growth of herbaceous plants in a natural vegetation mosaic in Niger. *J. Arid Environ.*
 1561 36,87-102
- 1562 Schlesinger, W.H., Peterjohn, W.T., 1991. Processes controlling ammonia volatilization from
 1563 Chihuahuan desert soils. *Soil Biology and Biochemistry* 23, 637-642.
- 1564 Schlesinger, W.H., Reynolds, J.F., Cunningham, G.L., Huenneke, L.F., Jarrell, W.M., Virginia,
 1565 R.A., Whitford, W.G., 1990. Biological feedbacks in global desertification. *Science*
 1566 247,1043-1048.
- 1567 Schlesinger, W.H., Raikes, J.A., Hartley, A.E., Cross, A.F., 1996. On the spatial pattern of soil
 1568 nutrients in desert ecosystems. *Ecology* 77,364-374.
- 1569 Scoging, H. 1992. Modelling overland-flow hydrology for dynamic hydraulics. In *Overland*
 1570 *Flow Hydraulics and Erosion Mechanics*. A. J. Parsons and A. D. Abrahams (eds).
 1571 London, UCL Press Limited.
- 1572 Scott, R.L., Cable, W.L., Hultine, K.R., 2008. The ecohydrological significance of hydraulic
 1573 redistribution in a semiarid savanna. *Water Resour. Res.* 44
- 1574 Smith, G., Holechek, J.L., Cardenas, M., 1996. Wildlife numbers on excellent and good
 1575 condition Chihuahuan Desert rangelands: an observation. *J. Range Manage.* 49, 489-493.

- 1576 Slatyer, R.O., 1961. Methodology of a water balance study conducted on a desert woodland
 1577 (*Acacia anuera* F. Muell) community in central Australia. UNESCO Arid Zone Res.
 1578 16,15-26.
- 1579 Starfield, A.M., 1996. Model of transient changes in arctic and boreal vegetation in response to
 1580 climate and land use change. *Ecological Applications* 6,842-864.
- 1581 Sun, G., Coffin, D.P., Lauenroth, W.K., 1998. Comparison of root distributions of species of
 1582 North American grasslands using GIS. *J. Vegetation Science*. 8, 587-596.
- 1583 Taylor, P.D., Fahrig, L., Henein, K., Merriam, G., 1993. Connectivity is a vital element of
 1584 landscape structure. *Oikos* 68, 571–573.
- 1585 Thornes, J.B., 1990. The interaction of erosional and vegetational dynamics in land degradation:
 1586 spatial outcomes. In JB Thornes (ed.) *Vegetation and Erosion*, John Wiley and Sons,
 1587 Chichester, pp 41–53
- 1588 Thornes, J. B., 2007. Modelling Soil Erosion by Grazing: Recent Developments
 1589 and New Approaches. *Geographical Research* 45: 13-26
- 1590 Thornes, J.B., Brandt, J. 1993. Erosion-Vegetation competition in a stochastic environment
 1591 undergoing climatic change. In AC Millington and KJ Pye, (eds.) *Environmental Change*
 1592 *in the Drylands: Biogeographical and Geomorphological Responses*, John Wiley and Sons,
 1593 Chichester, pp 306–320
- 1594 Tongway, D.J., Ludwig, J.A. (Editors), 2001. *Theories on the origins, maintenance, dynamics*
 1595 *and functioning of banded landscapes*. Springer, Heidelberg.
- 1596 Turchin, P., 2003. *Complex Population Dynamics*. Princeton University Press, Princeton.
- 1597 Turing, A.M., 1952. The chemical basis of morphogenesis. *Phil. Trans. Royal Soc. London B*
 1598 *Bio. Sci.* 237,37-72.

- 1599 Turnbull, L., Wainwright, J. Brazier, R.E., 2008. A conceptual framework for understanding
 1600 semi-arid land degradation: ecohydrological interactions across multiple-space and time
 1601 scales. *Ecohydrology* 1, 23–34.
- 1602 Turnbull, L., Wainwright, J., Brazier, R.E., 2010. Nitrogen and phosphorus dynamics during
 1603 runoff events over a transition from grassland to shrubland in the south-western United
 1604 States. *Hydrological Processes* 24, 393–414. DOI: 10.1002/hyp.7806
- 1605 Turner, M.G., Romme, W.H., Gardner, R.H., O’Neill, R.V., Kratz, T.K., 1993. A revised
 1606 concept of landscape equilibrium: disturbance and stability on scaled landscapes.
 1607 *Landscape Ecology* 8, 213–227.
- 1608 van de Koppel, J., Rietkerk, M., 2004. Spatial interactions and resilience in arid ecosystems. *Am.*
 1609 *Naturalist* 163,113-121.
- 1610 Versteeg, H.K., Malalasekera, W., 1995. *An Introduction to Computational Fluid Dynamics: The*
 1611 *Finite Volume Method*. Prentice Hall, London.
- 1612 Wainwright, J., 2009. Desert ecogeomorphology, in AJ Parsons and AD Abrahams (eds)
 1613 *Geomorphology of Desert Environments*, 2nd edition, Springer, Berlin.
- 1614 Wainwright, J., Bracken, L.J., 2011. Overland flow and runoff generation, in DSG Thomas (ed.)
 1615 *Arid Zone Geomorphology*, 3rd ed., 235–268, John Wiley and Sons, Chichester.
- 1616 Wainwright, J. and Parsons, A.J. 2010 Thornes, J.B. 1985: The ecology of erosion. *Geography* 70,
 1617 222-35. *Progress in Physical Geography*, 34, 399-408.
- 1618 Wainwright, J., Parsons, A.J., Abrahams, A.D., 1999. Field and computer simulation
 1619 experiments on the formation of desert pavement. *Earth Surf. Process. Landforms*. 24,
 1620 1025-1037.

- 1621 Wainwright, J., Parsons, A.J., Abrahams, A.D., 2000. Plot-scale studies of vegetation, overland
 1622 flow and erosion interactions: case studies from Arizona and New Mexico. *Hydrol. Proc.*
 1623 14,2921–2943.
- 1624 Wainwright, J., Parsons, A.J., Schlesinger, W.H., Abrahams, A.D., 2002. Hydrology–vegetation
 1625 interactions in areas of discontinuous flow on a semi-arid bajada, southern New Mexico. *J.*
 1626 *Arid Environments* 51, 319-330.
- 1627 Wainwright, J., Parsons, A.J., 2002. The effect of temporal variations in rainfall on scale
 1628 dependency in runoff coefficients. *Water Resour. Res.* 38(12), 1271.
- 1629 Wainwright, J., 2005. Climate and Climatological Variations in the Jornada Experimental Range
 1630 and Neighbouring Areas of the US Southwest. *Adv. Environ. Monitor. Model.* 1,39-110.
- 1631 Wainwright, J.W., Parsons, A.J., Müller, E.N., Brazier, R.E., Powell, D.M., Fenti, B., 2008a. A
 1632 Transport-Distance Approach to Scaling Erosion Rates: 1. Background and Model
 1633 Development. *Earth Surf. Process. Landforms.* 33, 813–826.
- 1634 Wainwright, J.W., Parsons, A.J., Müller, E.N., Brazier, R.E., Powell, D.M., Fenti, B. 2008b. A
 1635 Transport-Distance Approach to Scaling Erosion Rates: 2. Sensitivity and Evaluation of
 1636 MAHLERAN. *Earth Surf. Process. Landforms.* 33, 962–984.
- 1637 Wainwright, J.W., Parsons, A.J., Müller, E.N., Brazier, R.E., Powell, D.M., Fenti, B., 2008c. A
 1638 Transport-Distance Approach to Scaling Erosion Rates: 3. Evaluating Scaling
 1639 Characteristics of MAHLERAN. *Earth Surf. Process. Landforms.* 33, 1113–1128.
- 1640 Wainwright, J. 2009. Desert ecogeomorphology. In AJ Parsons and AD Abrahams (eds)
 1641 *Geomorphology of Desert Environments*, Springer, Berlin
- 1642 Walter, H., 1971. *Ecology of Tropical and Subtropical Vegetation*. Oliver and Boyd, Edinburgh.
- 1643 Ward, D 2008 *The Biology of Deserts*. Oxford University Press, Oxford.

- 1644 Went, F.W., Westergaard, M., 1949. Ecology of desert plants. III. Development of plants in the
 1645 Death Valley National Monument, California. *Ecology*. 30, 26-38.
- 1646 Western AW, Bloschl G, and Grayson RB. 2001. Towards capturing hydrologically significant
 1647 connectivity in spatial patterns. *Water Resources Research* **37**, 83–97.
- 1648 Westoby, M., 1980. Elements of a theory of vegetation dynamics in arid rangelands. *Israel J.*
 1649 *Botany* 28,169-194
- 1650 Westoby, M., Walker, B., Noy-Meir, I., 1989. Opportunistic management for rangelands not at
 1651 equilibrium. *J. Range Manage.* 42(4),266-274
- 1652 White, L.P., 1997. Vegetation stripes on sheet wash surfaces. *J. Ecology* 59, 615-622
- 1653 Whitford, W.G., 2002. *Ecology of Desert Systems*, San Diego, Academic Press, 343pp.
- 1654 Williams, C.A., Albertson, J.D., 2006. Dynamical effects of the statistical structure of annual
 1655 rainfall on dryland vegetation. *Global Change Biology*. 12, 777-792.
- 1656 Yao, J., Peters, D.P.C., Havstad, K.M., Gibbens, R.P., Herrick, J.E., 2006. Multi-scale factors
 1657 and long-term responses of Chihuahuan Desert grasses to drought. *Landscape Ecology*
 1658 21,1217-1231.
- 1659

1660 **ECOLOGICAL ARCHIVES MATERIAL**

1661 Appendix: Exploration of Model Behaviour

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Table 1 Summary of input conditions used in model simulations

Length of rainfall record used in simulations	312 years, 80 years, or 80-year average
Size of grid	50 x 50 m
Gradient of slope	2 degrees
Nitrogen input	0.65g m ⁻²
Water input to each cell	Constant, long term average = 228 mm yr ⁻¹ Stochastic, average = 243 mm yr ⁻¹ Instrumented, average = 243 mm yr ⁻¹

Table 2 Ranges of biomass consumption under three different grazing strategies
(Havstad *et al.*, 2006)

Consumption of palatable biomass in $\text{g m}^{-2} \text{yr}^{-1}$ and converted to a percentage of biomass consumed as applied in the model			
Range	Conservative	Recommended	Overstocked
Lowest quoted value	$8 \text{ g/m}^{-2} \text{yr}^{-1}$ (2.5%)	$7 \text{ g/m}^{-2} \text{yr}^{-1}$ (2.2%)	$30 \text{ g/m}^{-2} \text{yr}^{-1}$ (9.4%)
Highest quoted value	$14 \text{ g/m}^{-2} \text{yr}^{-1}$ (4.4%)	$21 \text{ g/m}^{-2} \text{yr}^{-1}$ (6.6%)	$60 \text{ g/m}^{-2} \text{yr}^{-1}$ (18.8%)

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UP SLOPE (NORTH)				
EAST	0.0	0.1	0.0	WEST
	0.1	0.5	0.1	
	0.05	0.1	0.05	
DOWN SLOPE (SOUTH)				
Percentage of resource moved out from a cell containing black grama by the water vector				

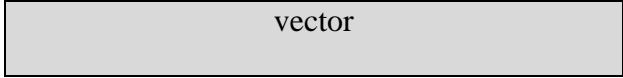
UP SLOPE (NORTH)				
EAST	0.0	0.2	0.0	WEST
	0.05	0.5	0.05	
	0.05	0.1	0.05	
DOWN SLOPE (SOUTH)				
Percentage of resource moved out from a cell containing creosotebush by the water vector				

UP SLOPE				
EAST	0.0	0.05	0.0	WEST
	0.15	0.5	0.15	
	0.05	0.05	0.05	
DOWN SLOPE				
Percentage of nitrogen moved out from a cell containing black grama by the wind vector				

UP SLOPE				
EAST	0.0	0.05	0.0	WEST
	0.05	0.45	0.05	
	0.05	0.3	0.05	
DOWN SLOPE				
Percentage of nitrogen moved out from a cell containing creosotebush by the wind vector				

UP SLOPE (SOUTH)				
EAST	0.1	0.1	0.1	WEST
	0.1	0.2	0.1	
	0.1	0.1	0.1	
DOWN SLOPE (NORTH)				
Percentage of nitrogen moved out from a cell containing black grama by the grazing vector				

UP SLOPE (SOUTH)				
EAST	0.0625	0.0625	0.0625	WEST
	0.0625	0.5	0.0625	
	0.0625	0.0625	0.0625	
DOWN SLOPE (NORTH)				
Percentage of nitrogen moved out from a cell containing creosotebush by the grazing				



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Table 4 Summary of species demographic data

Demographic data	Grass	Shrub	Data source
Maximum annual growth rate (%)			Peters, (2002a)
Maximum biomass in 1-m ² cell (g)	319	222	Maneta <i>et al</i> (2008)
Water Efficiency (g water per g of biomass)	3.5	2.48	Peters, (2002a)
Nitrogen use efficiency (g nitrogen per g of biomass)	0.6206	0.2767	Peters, (2002a)
Water maintenance requirement (g water per g of biomass)	0.7	0.496	Peters, (2002a)
Nitrogen maintenance requirement (g nitrogen per g of biomass)	0.125	0.055	Wainwright (2009)
Failure rate of species (% of species in cell)	5	5	
Mortality rate of species (% of species in cell)	10	10	

Table 5 Percentage of roots of the grass and shrub species that are distributed between the three soil layers

1666

	<i>Bouteloua eriopoda</i>	<i>Larrea tridentate</i>
Top Layer	13.3%	6.7%
Middle Layer	50.4%	32.0%
Deep Layer	36.3%	61.3%

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Table 6 Percentage of propagules for each species that is amenable to movement by the vectors

Propagules	Vector		
	Water	Wind	Grazer
Black grama	75%	20%	5%
Creosotebush	85%	10%	5%

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LIST OF FIGURES

- Figure 1 Conceptual framework for modelling dynamic relationships between elements of desert ecosystems.
- Figure 2 The mechanisms for resource and propagule redistribution based on cell connectivity. If a cell is disconnected (i.e. receives no input of resources and propagules from up-vector cells), or is at the end of a connected pathway, available resources and propagules entering that cell are redistributed according to rules given in Table 3. If the cell lies on a connected pathway, available resources and propagules are moved down-vector.
- Figure 3 Simulation grid showing an arbitrary cell and its Moore neighbourhood. Labelling of the neighbours denotes their direction with respect to the grid.
- Figure 4 Schematic diagram of model execution in which QV is input from vertical processes; QH is redistribution by lateral processes; R is resources; P is propagules; and θ , φ and ψ are three temporal scales.
- Figure 5 a) Rainfall data reconstructed from tree ring data at the Jornada Headquarters from 1659 to 1969. Measured values from the

instrumental record are plotted from 1915 to 1995 for comparison (Wainwright, 2005). NB $\text{g m}^{-2} \text{a}^{-1}$ are equivalent to mm a^{-1}

b) Comparison of reconstructed with real rainfall data over common years at the Jornada Headquarters (Wainwright, 2005).

Figure 6 Stocking levels of cattle (represented by animal unit equivalents) in the Jornada basin from 1915 to 2001. An animal unit is a mature, non lactating 1100 lbs beef cow consuming 26 lbs of forage per day.

Figure 7 Schematic diagram to show distribution of roots in relation to the three soil layers in the model.

Figure 8 The randomly generated initial conditions for the model runs.

Figure 9 a) Lognormal probability distribution (P) shown by the dotted line to approximate observed rainfall (shown by the solid line) at the Jornada LTER

b) Example of the resulting stochastic rainfall model.

Figure 10 Averages of a) biomass of grass and shrubs; b) water and nitrogen levels in the mid and deep soil layers; c) connectivity with respect to the water vector for the 50 cells along the centre line of the grid for each year of the

312-year stochastic rainfall simulation. Note that because the rainfall record is a stochastically generated one individual model realizations may differ substantially. However, all that we have run show no long-term spatial reorganisation of the vegetation.

Figure 11 Averages of a) biomass of grass and shrubs; b) water and nitrogen levels in the mid and deep soil layers; c) connectivity with respect to the water vector for the 50 cells along the centre line of the grid for each year of the 312-year reconstructed-rainfall series simulation.

Figure 12 Maps depicting the spatial distribution of grass and shrub biomass in years 1683, 1783, 1883, 1958 and 1970 for the reconstructed-rainfall series simulation.

Figure 13 Averages of i) biomass of grass and shrubs; ii) water and nitrogen levels in the mid and deep soil layers; iii) connectivity with respect to the water vector for the 50 cells along the centre line of the grid for each year of the 312-year reconstructed-rainfall series simulation, for a) a conservative grazing level, b) the recommended grazing level, and c) and the overgrazed case.

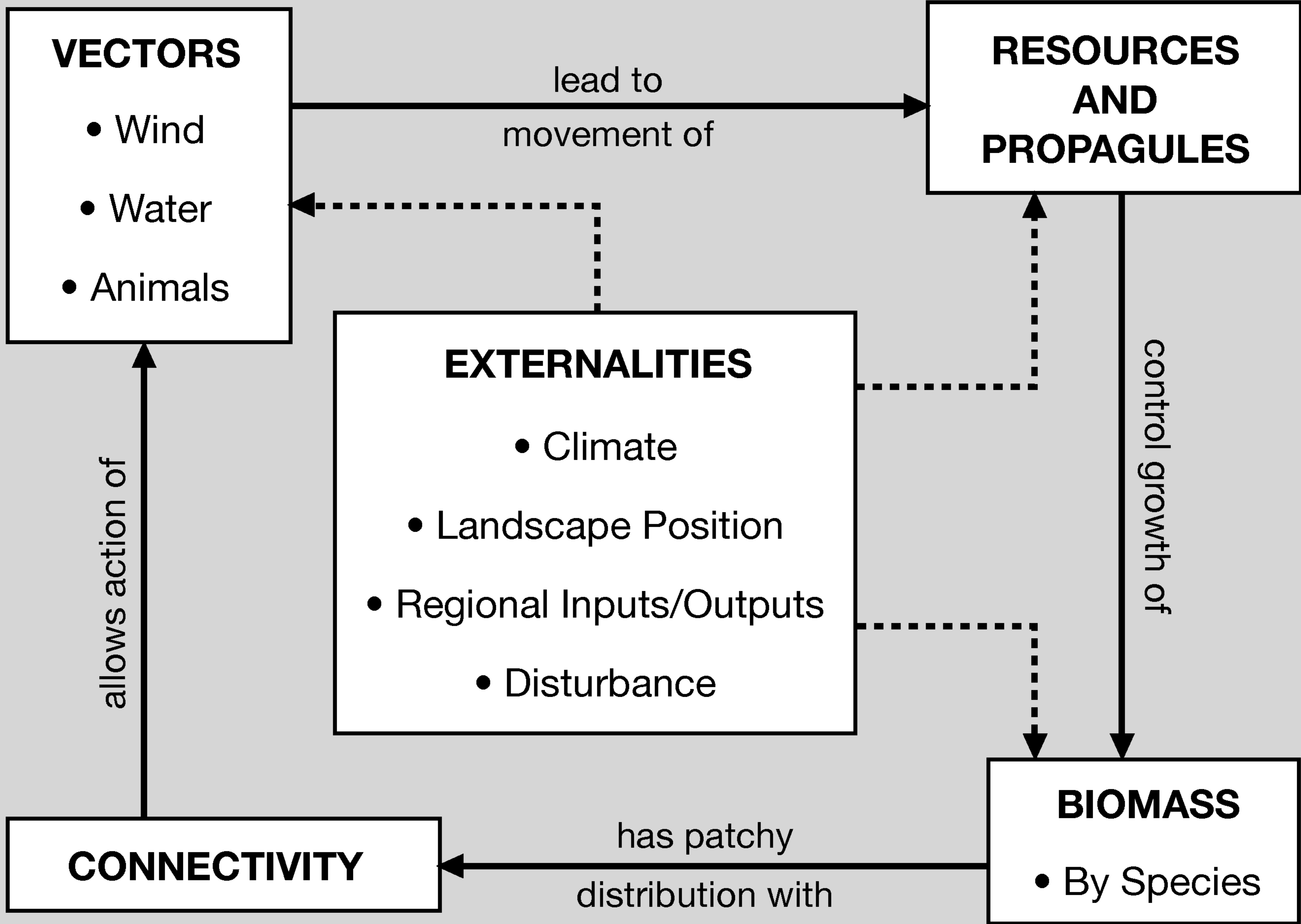
Figure 14 Maps depicting the spatial distribution of grass and shrub biomass in years 1683, 1783, 1883, 1958 and 1970 for the overgrazed -grazing simulation.

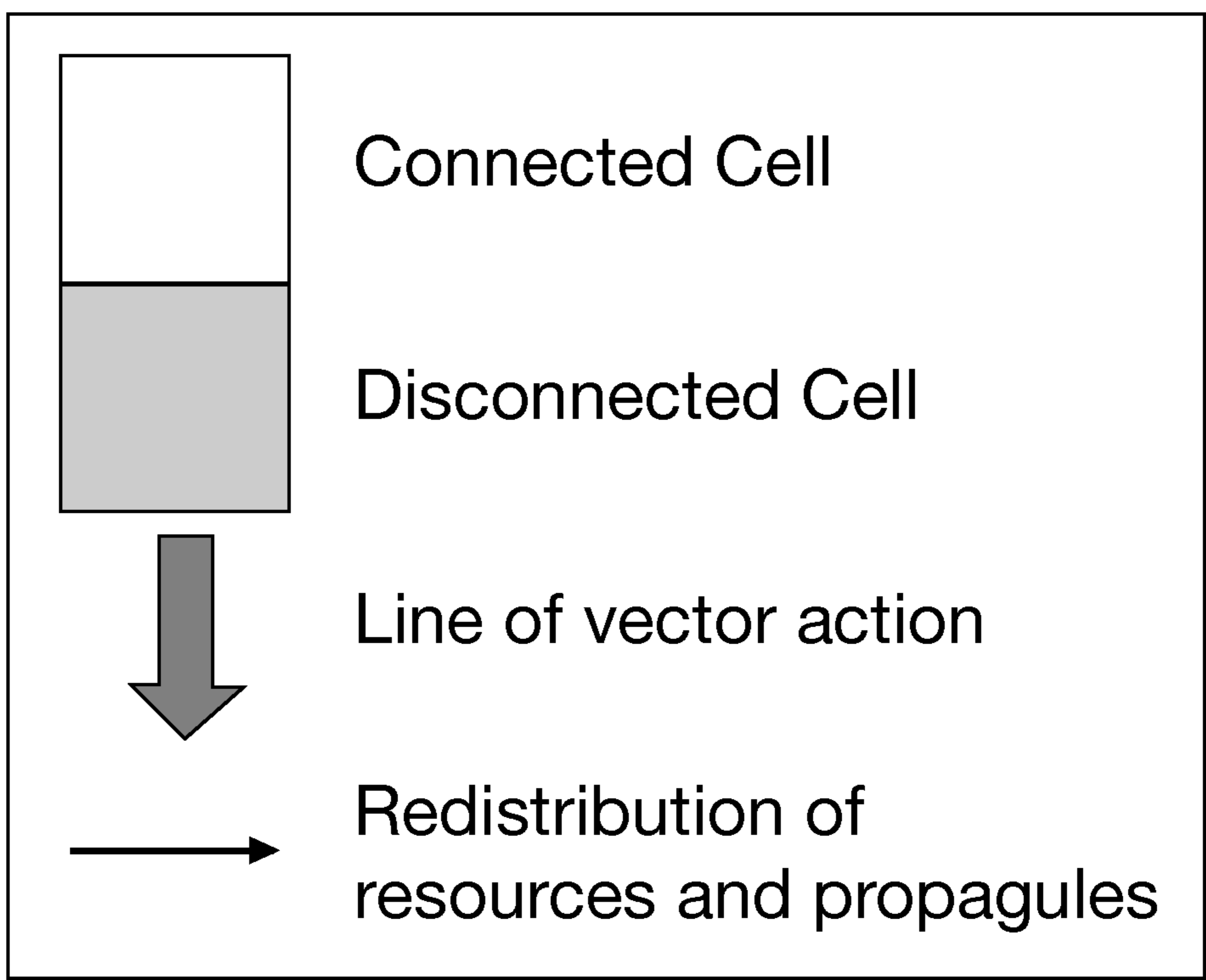
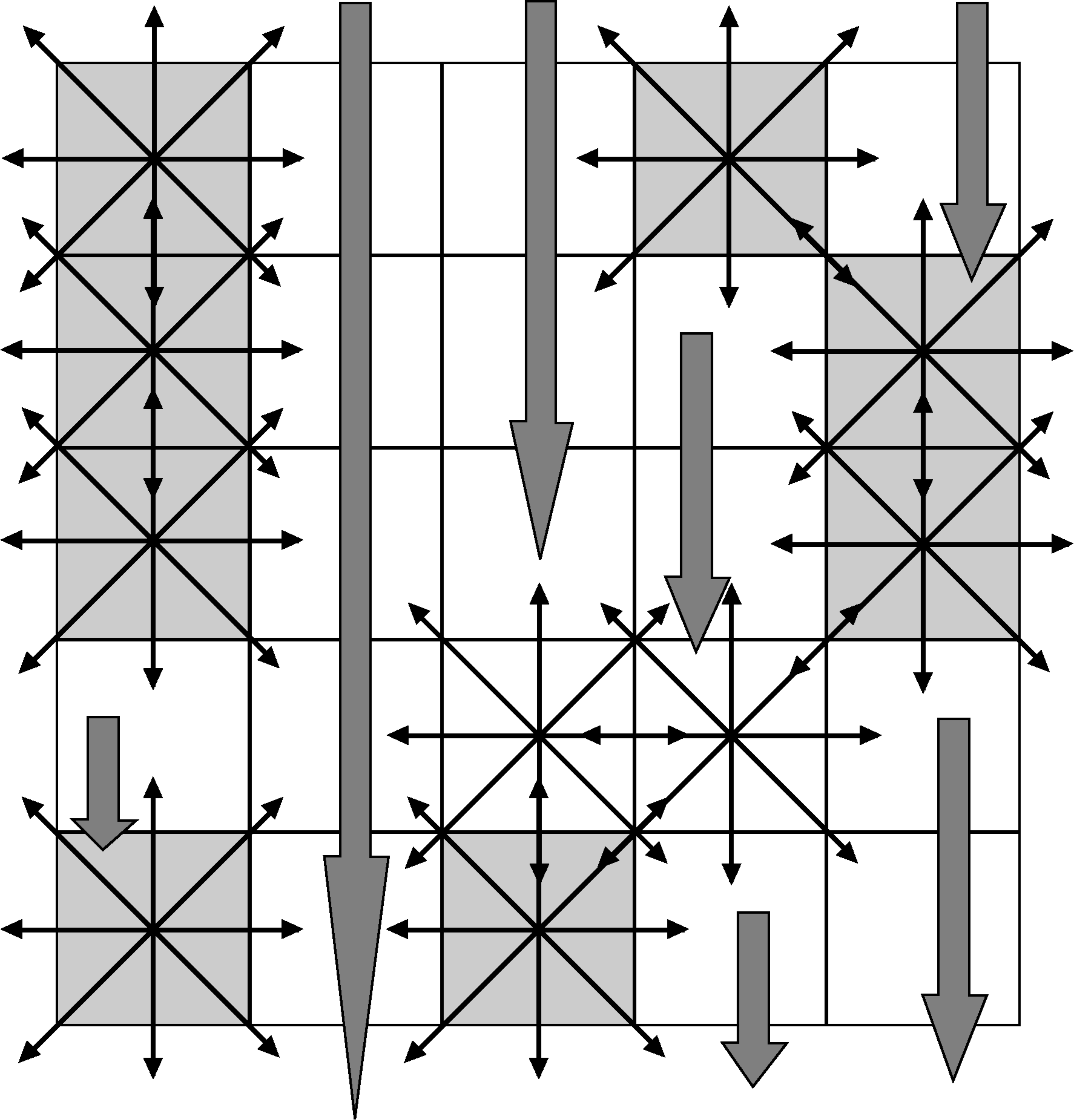
Figure 15 Averages of a) biomass of grass and shrubs; b) water and nitrogen levels in the mid and deep soil layers; c) connectivity with respect to the water vector for the 50 cells along the centre line of the grid for each year using the instrumented rainfall record and a conservative grazing level.

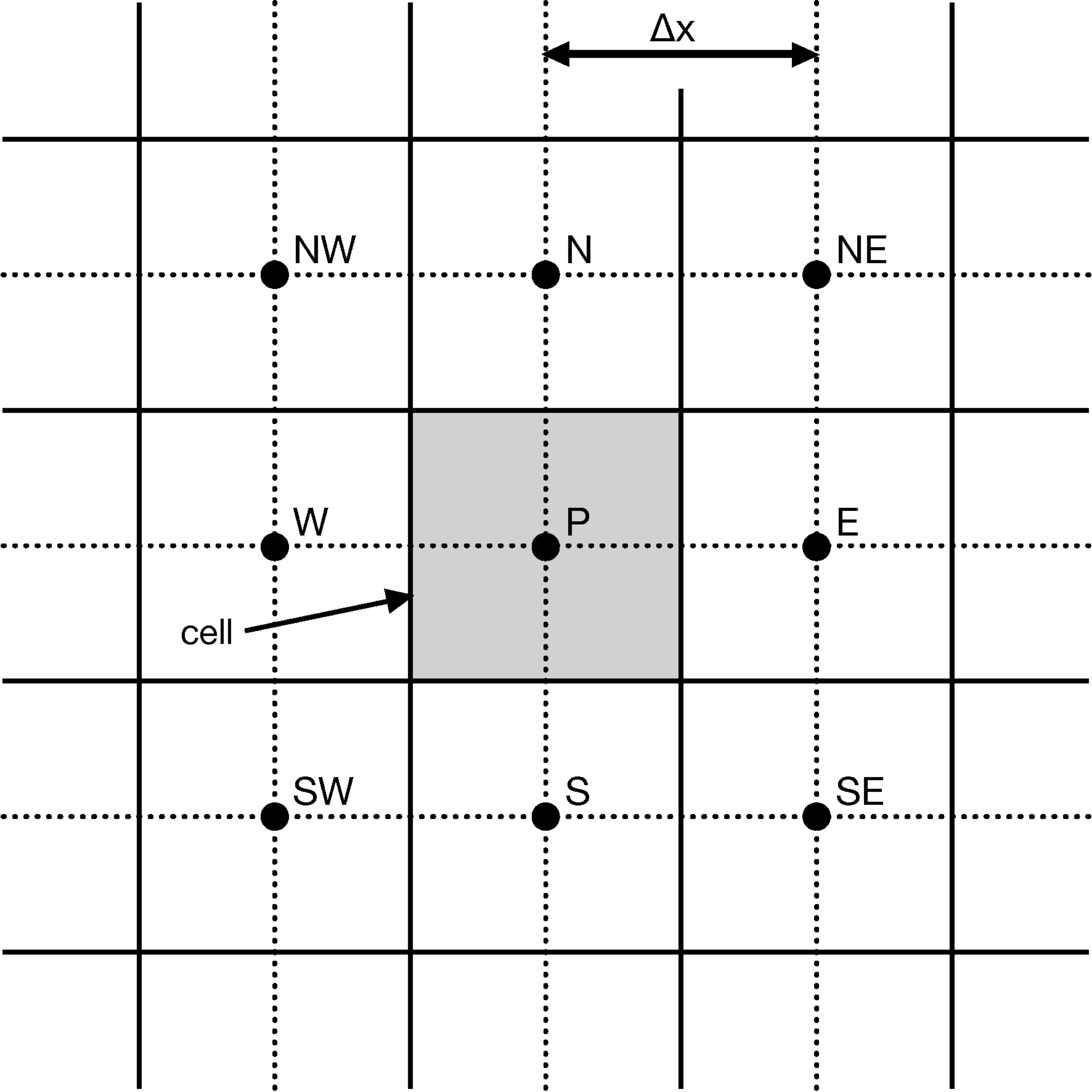
Figure 16 Maps depicting the spatial distribution of grass and shrub biomass at 10-year intervals for the simulation using the instrumented rainfall record and a conservative grazing level.

Figure 17 Biomass and resource distribution along the centre line of the grid at 10-year intervals for the simulation using the instrumented rainfall record and a conservative grazing level. For ease of representation, modelled biomass and resource density are scaled (normalised) the maximum potential biomass given in Peters (2002a).

Figure 18 Selected comparisons of individual cells taken from the centre line of the grid compared with experimental-plot data from Yao *et al* (2006). For ease of representation, modelled biomass is presented as a proportion (normalised biomass) of the maximum potential biomass given in Peters (2002a).







START

Define field size; species specific data;
number of species, vectors and
resources; simulation time

Time = end ?

Y

N

Each cell

Add QV values
for R and P to
each cell

$d\theta = d\phi ?$

Y

N

Each cell

Assign a proportion of
QV to each vector

Redistribute available
resources as QH

$d\theta = d\psi ?$

N

Y

Each cell

Change of biomass due
to new resource in cell

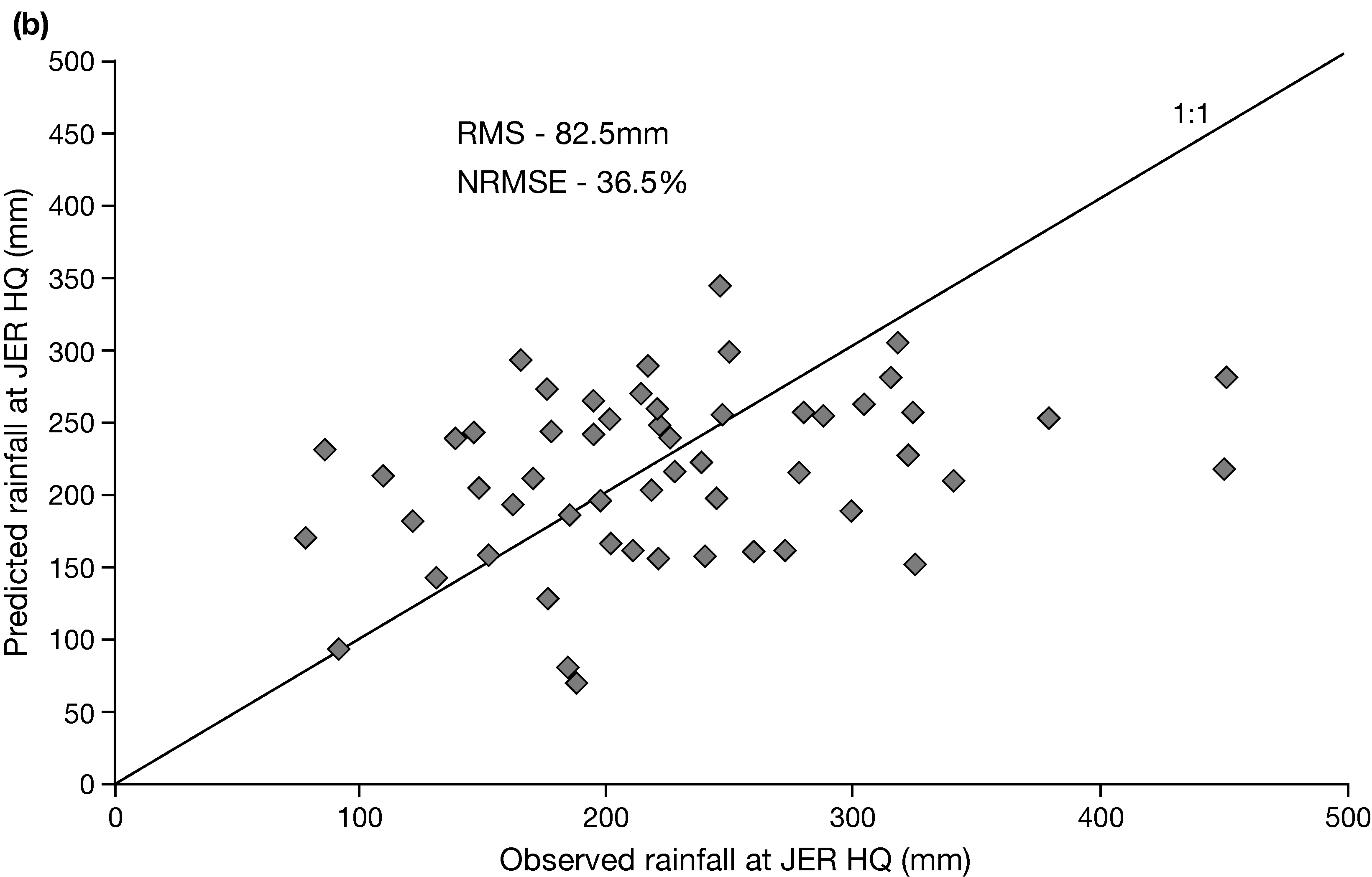
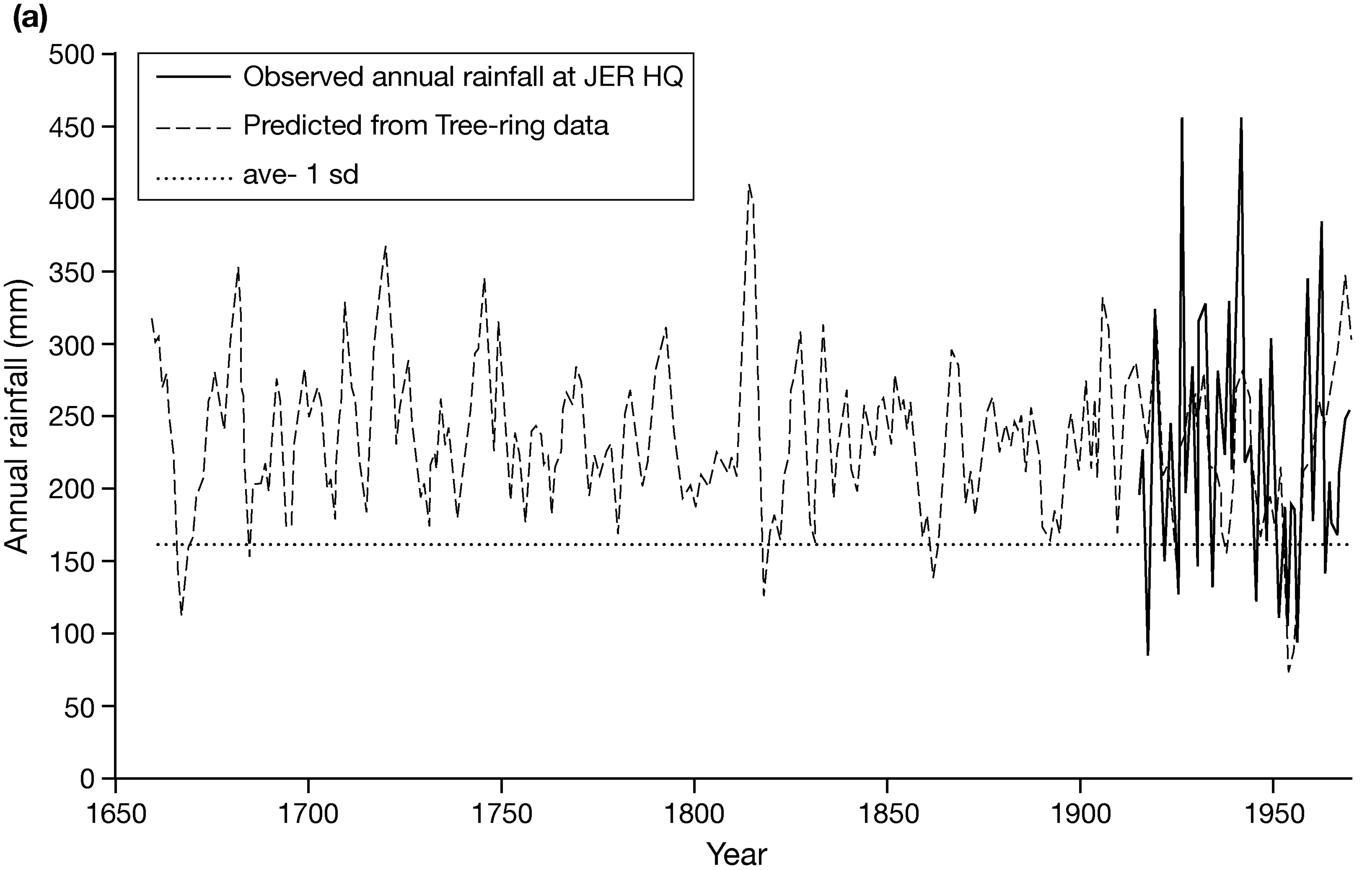
Calculate growth rate
in proportion to most
limiting resource

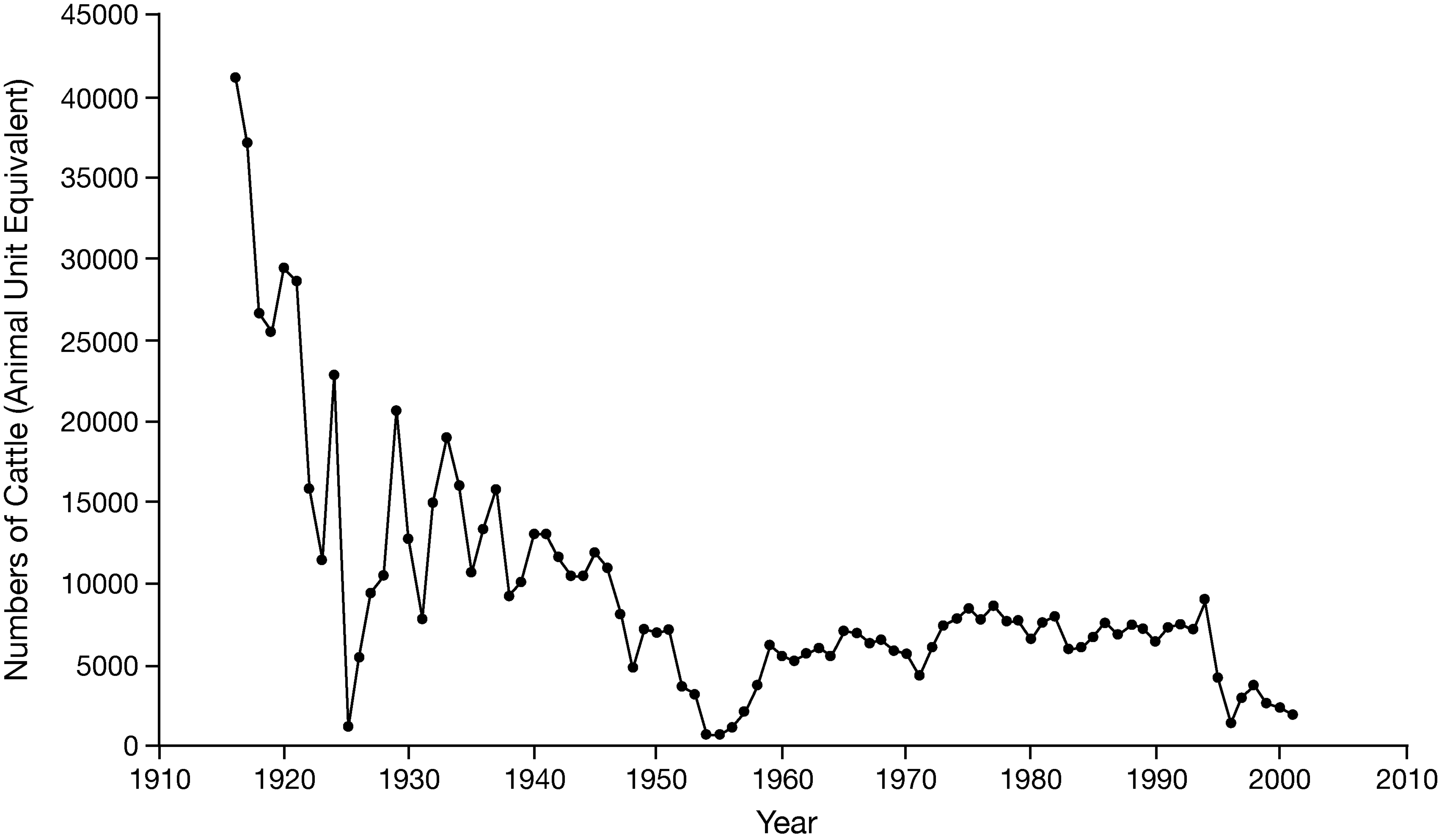
$d\psi = dt ?$

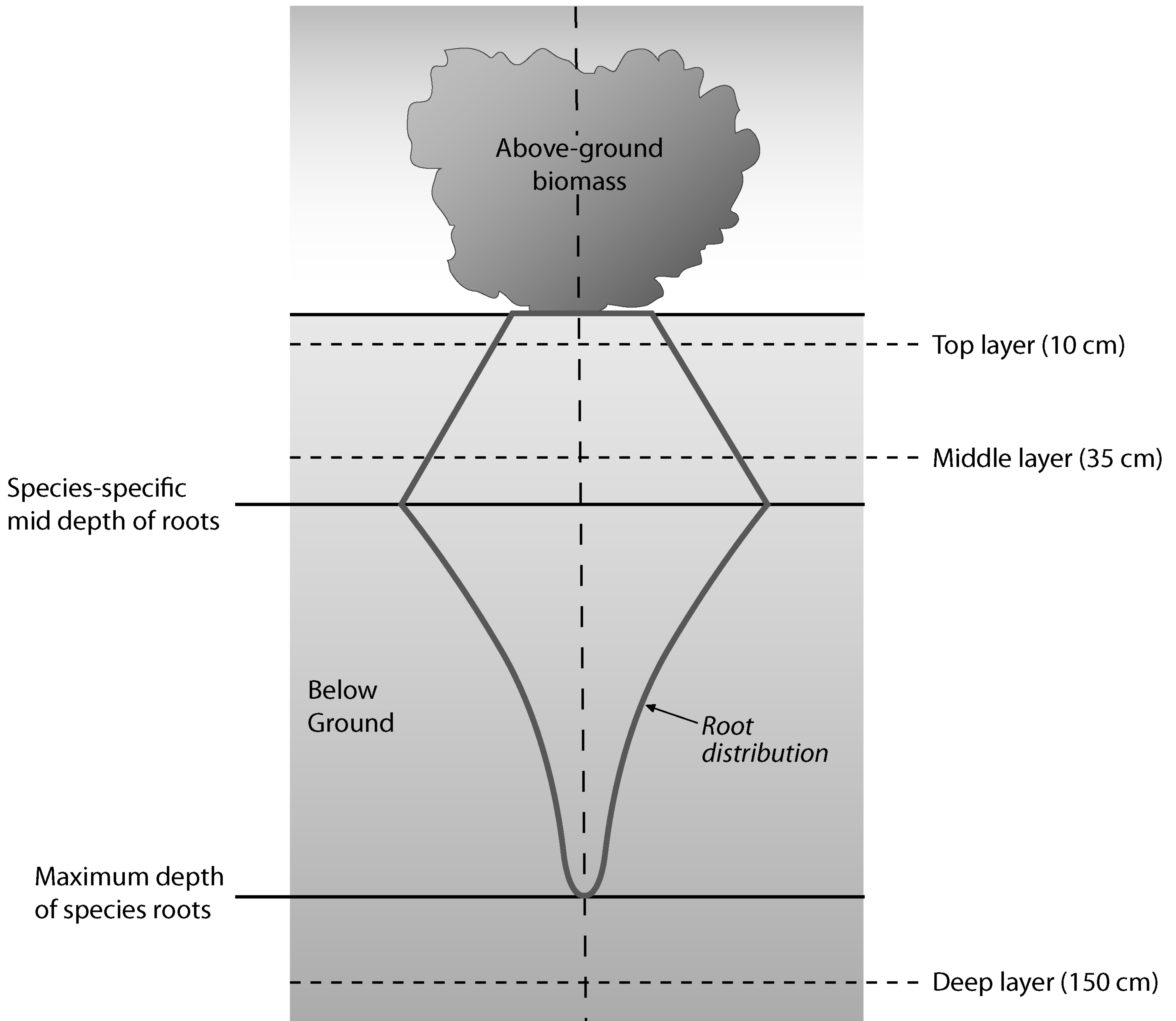
N

Y

END

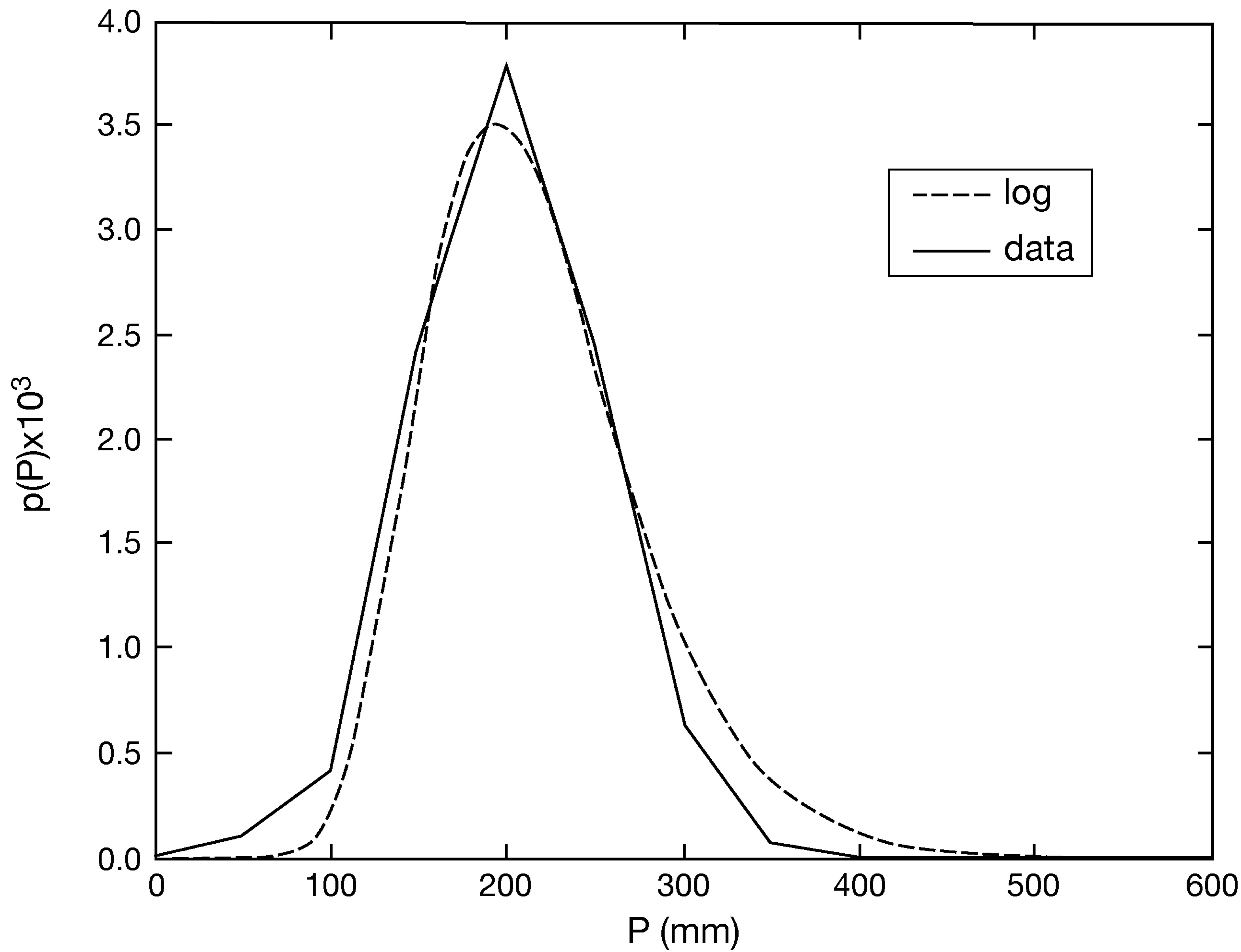
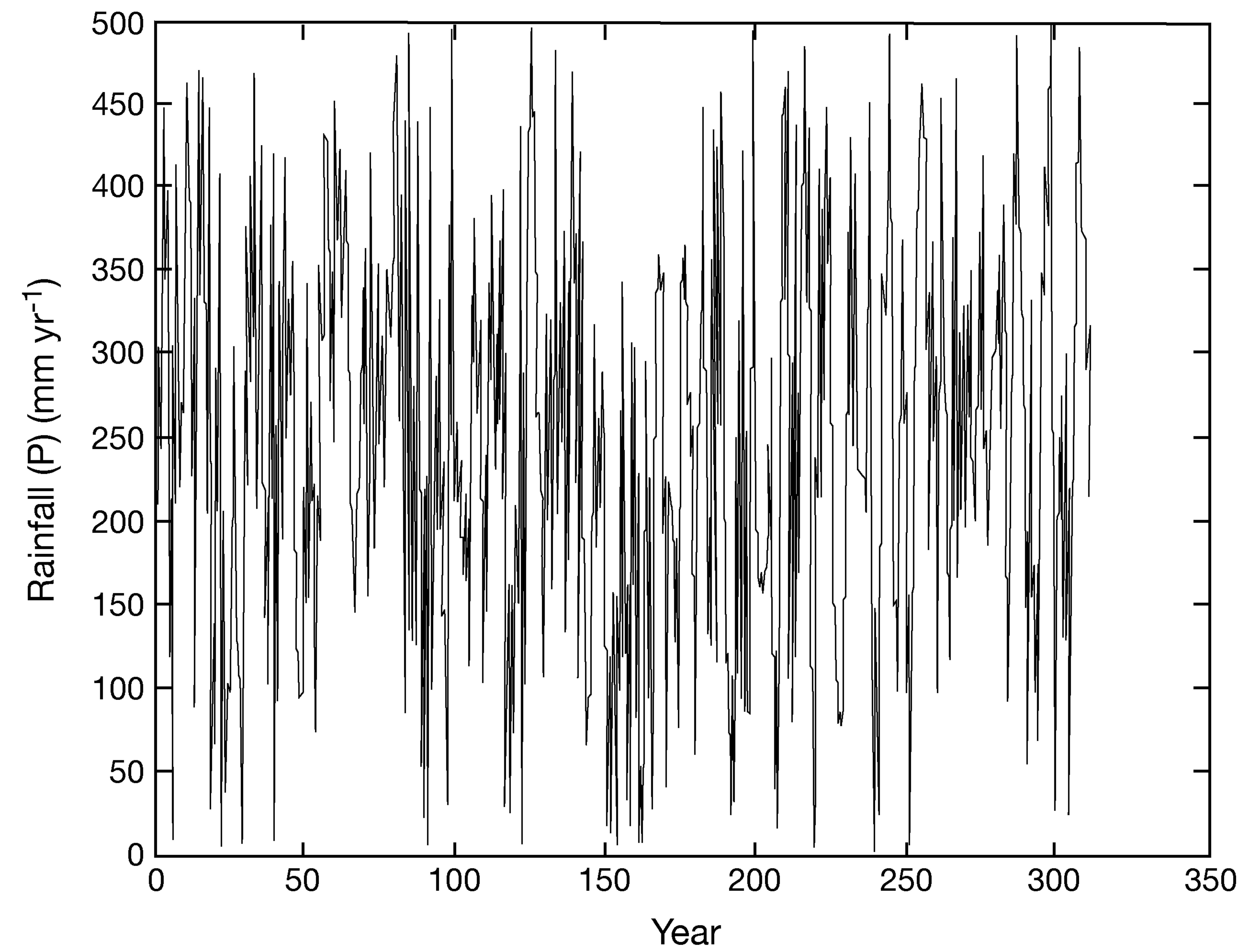


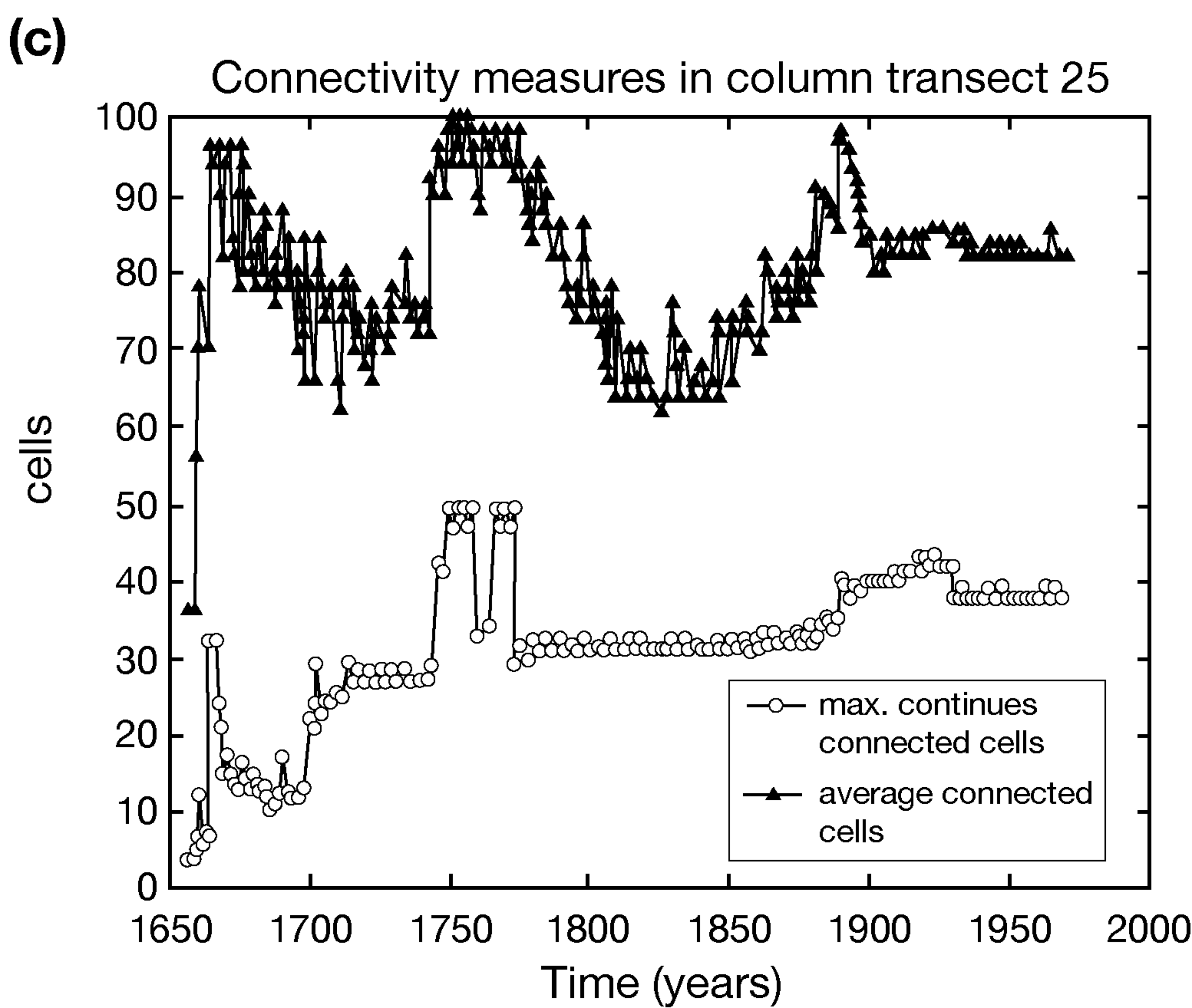
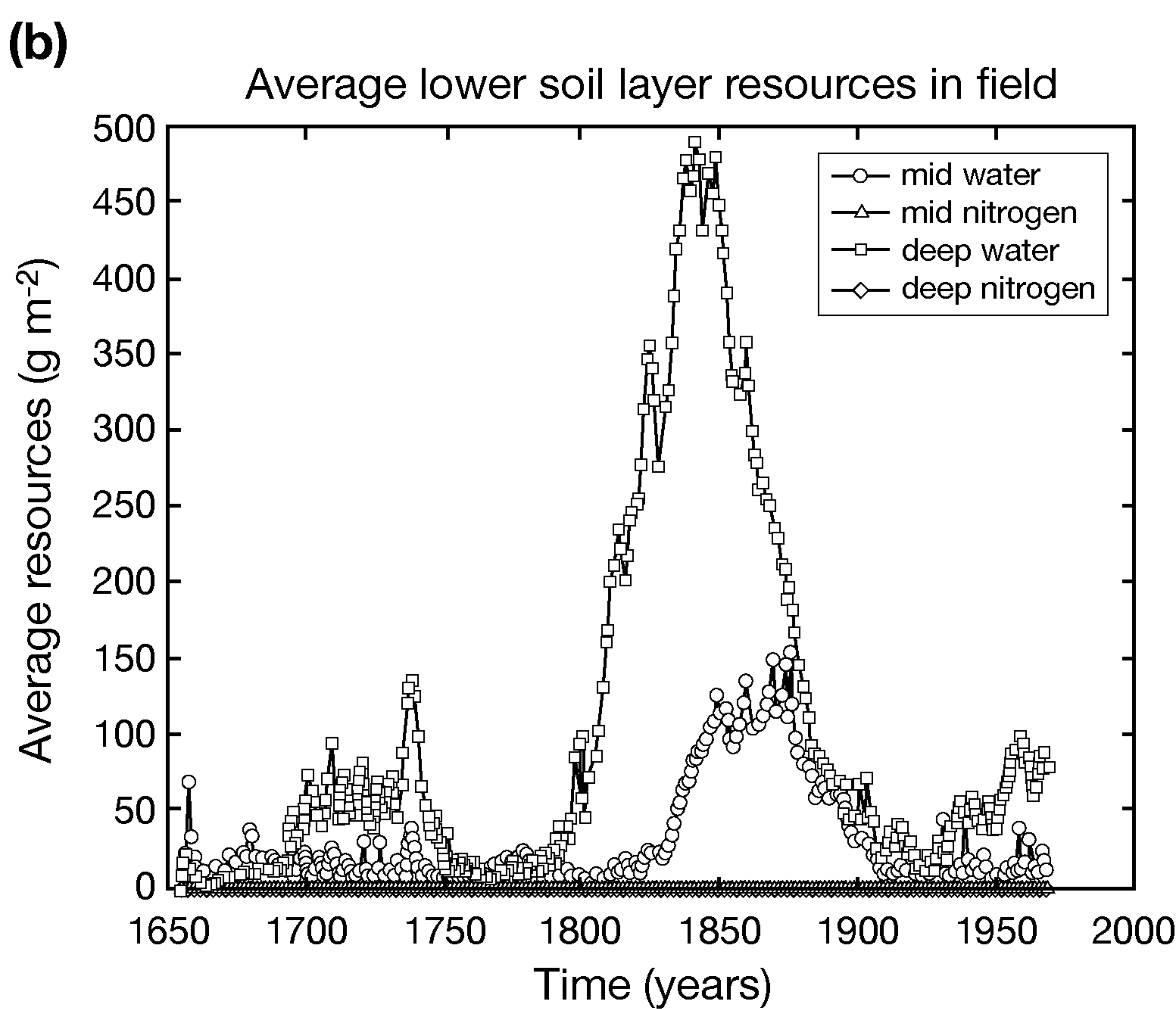
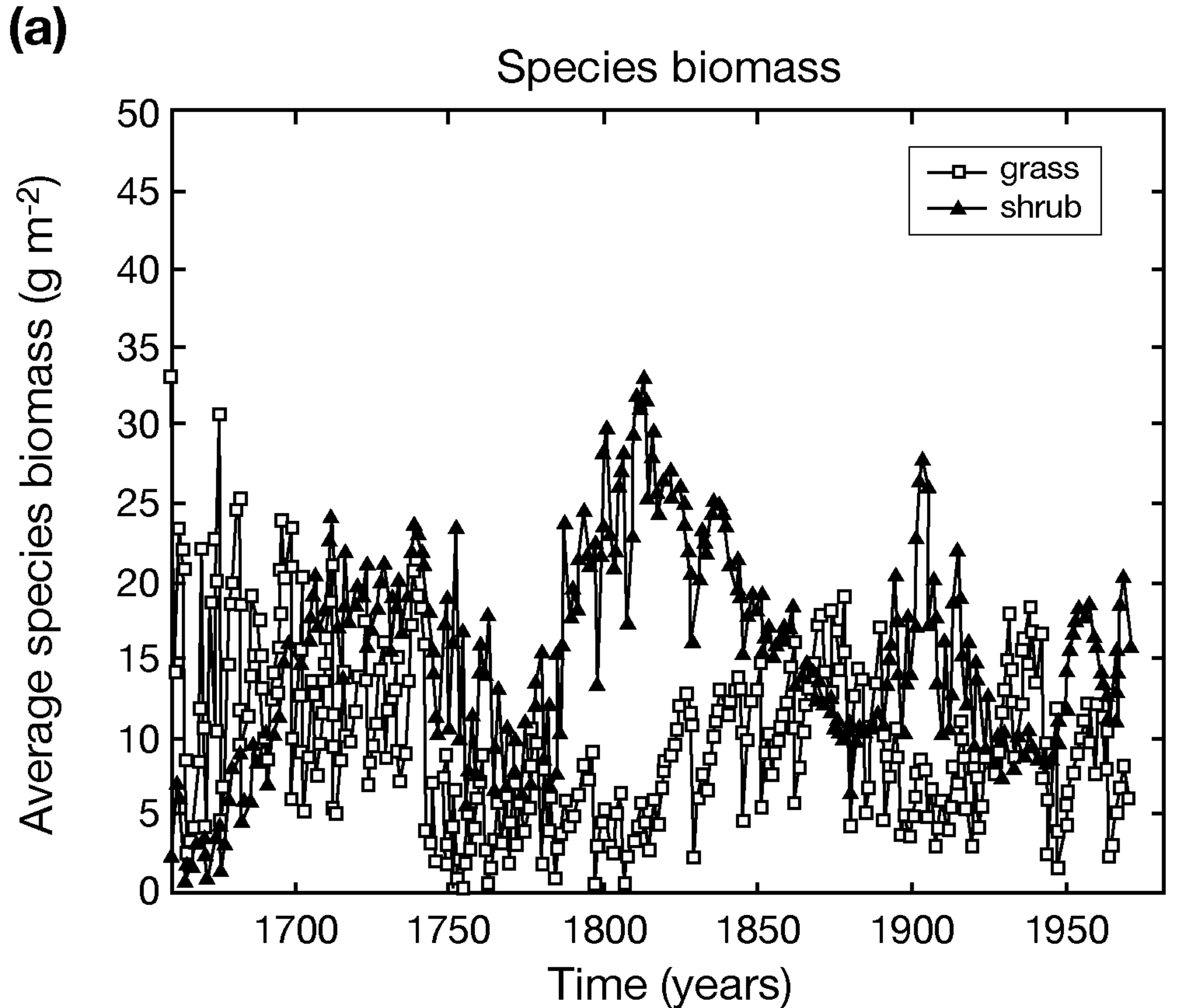


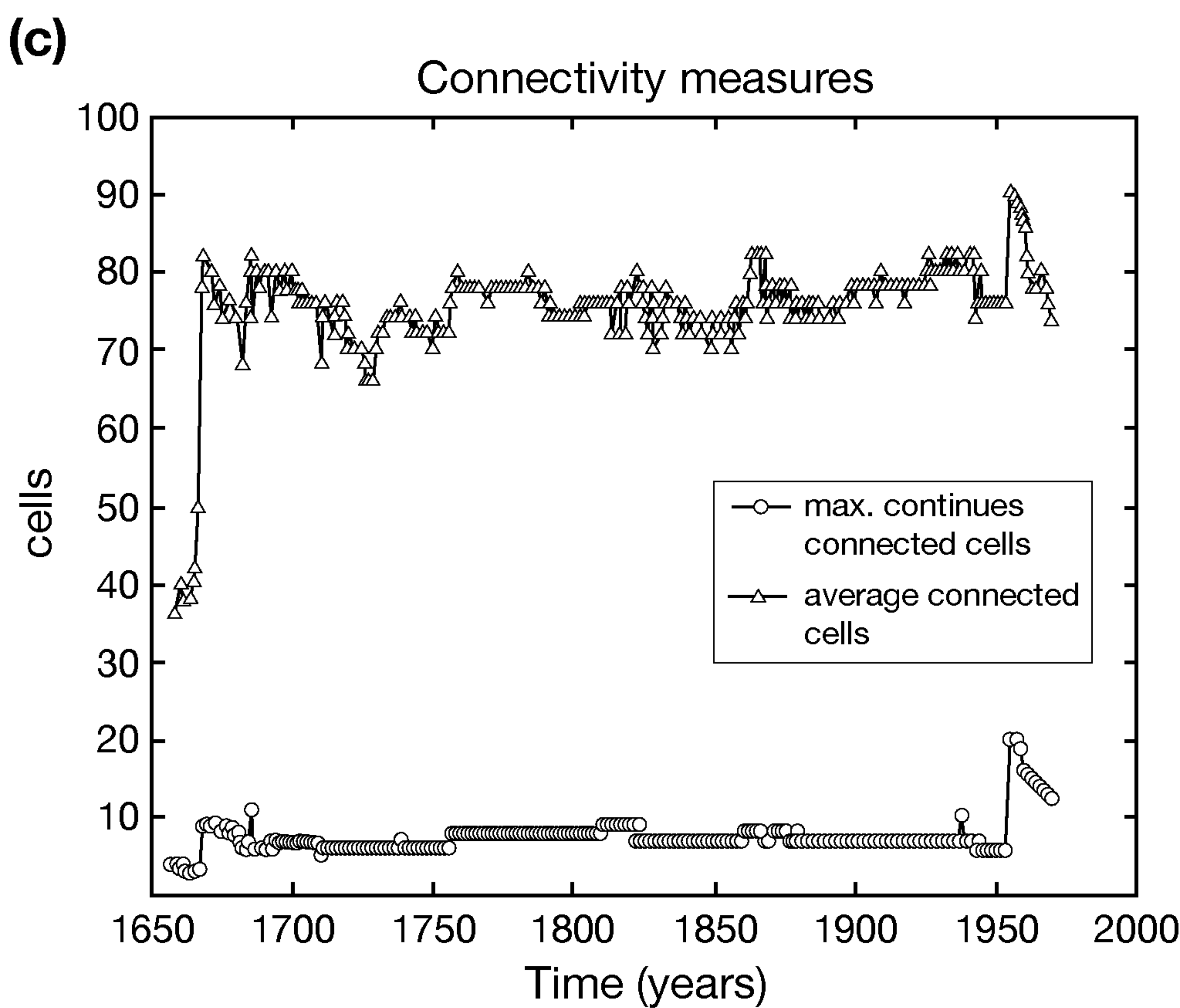
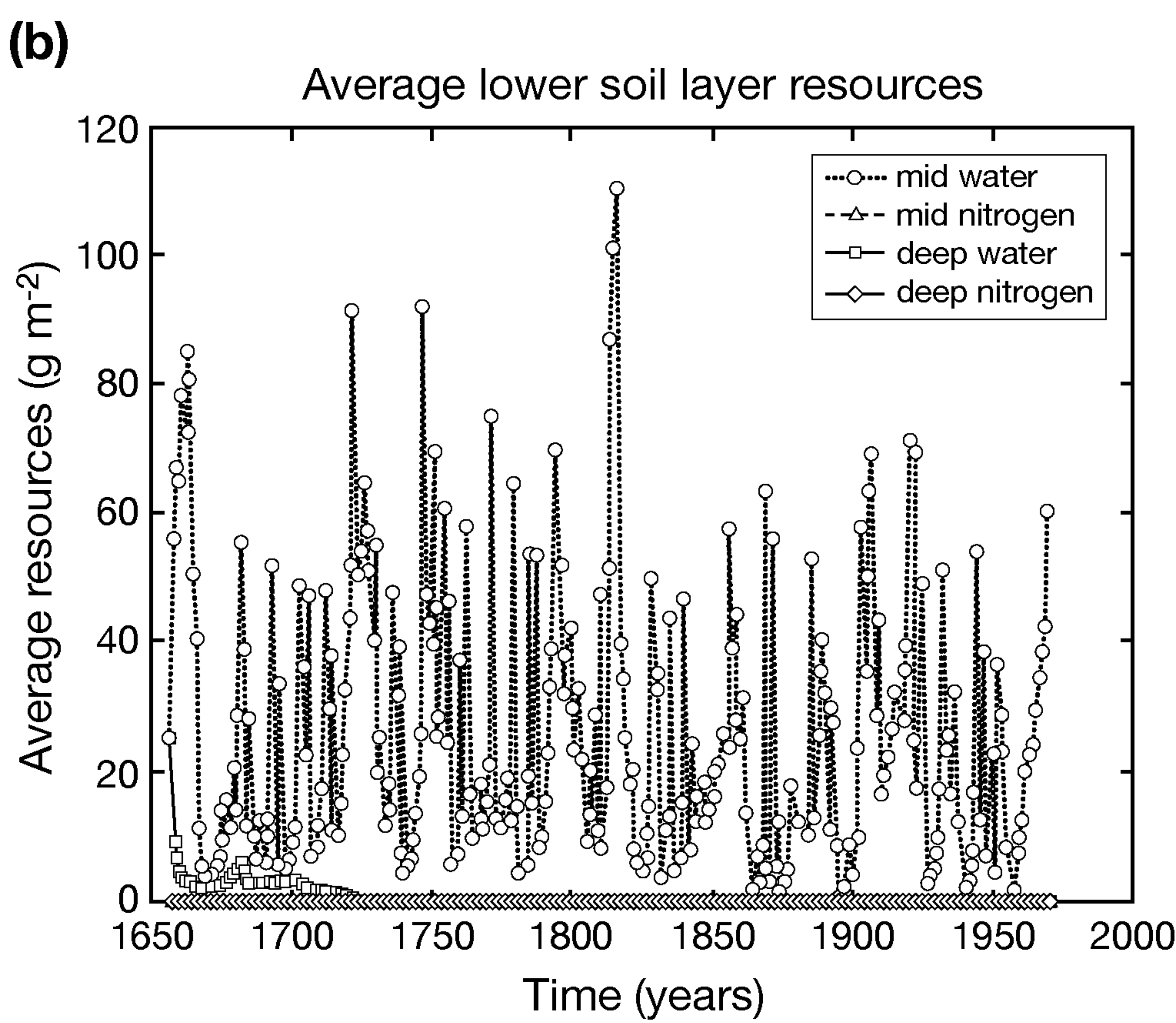
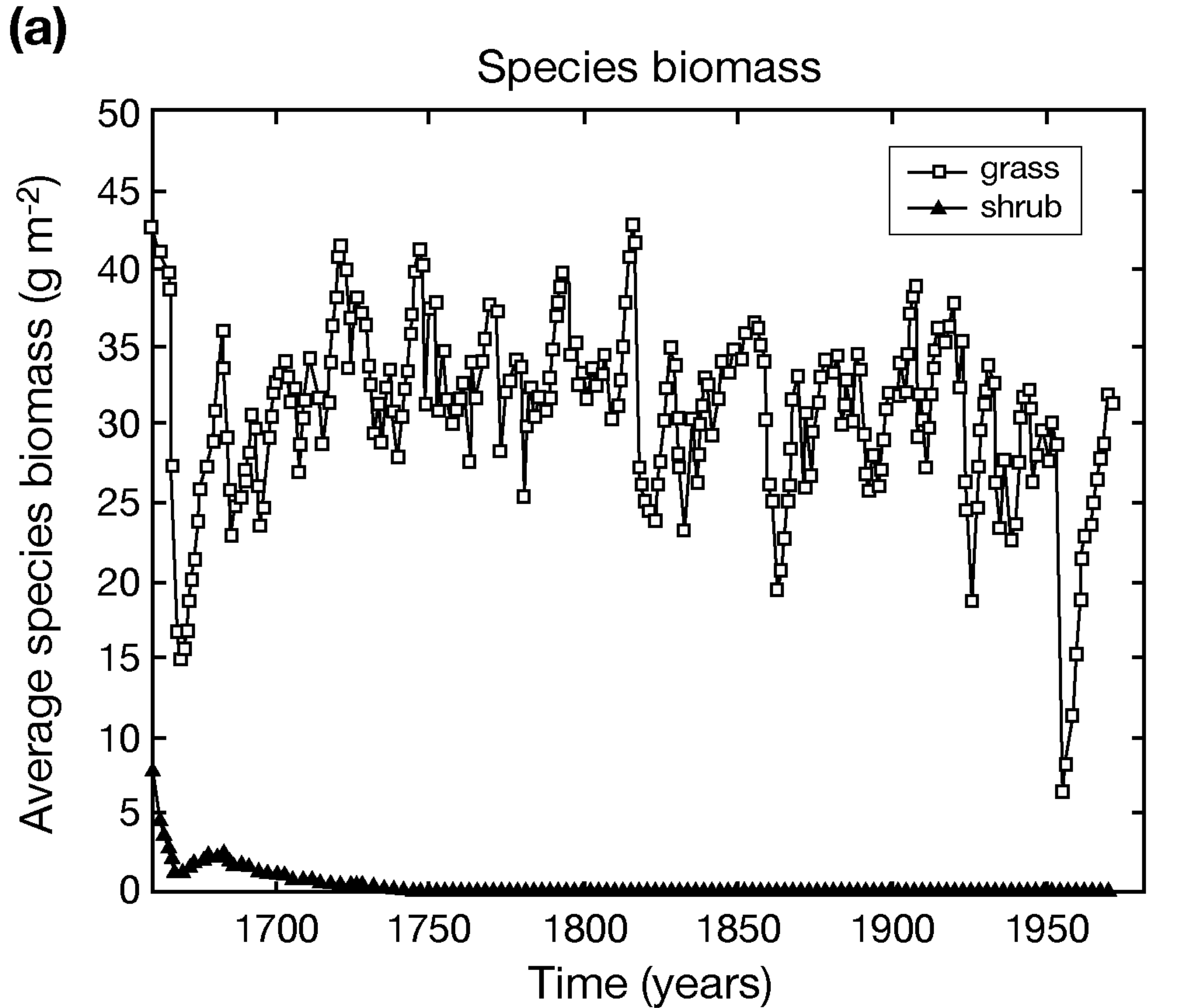


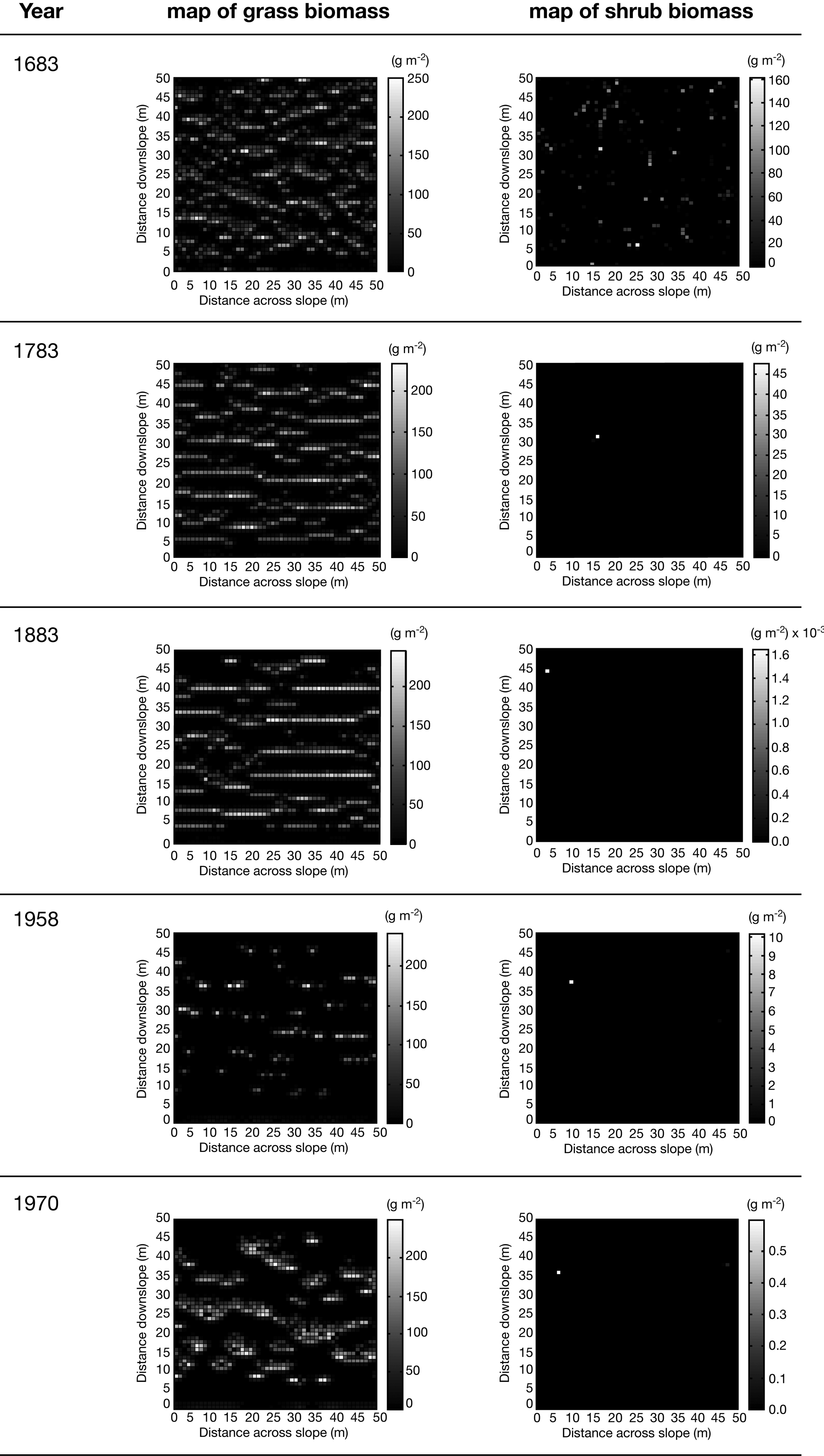
Initial distribution of grass biomass



(a)**(b)**





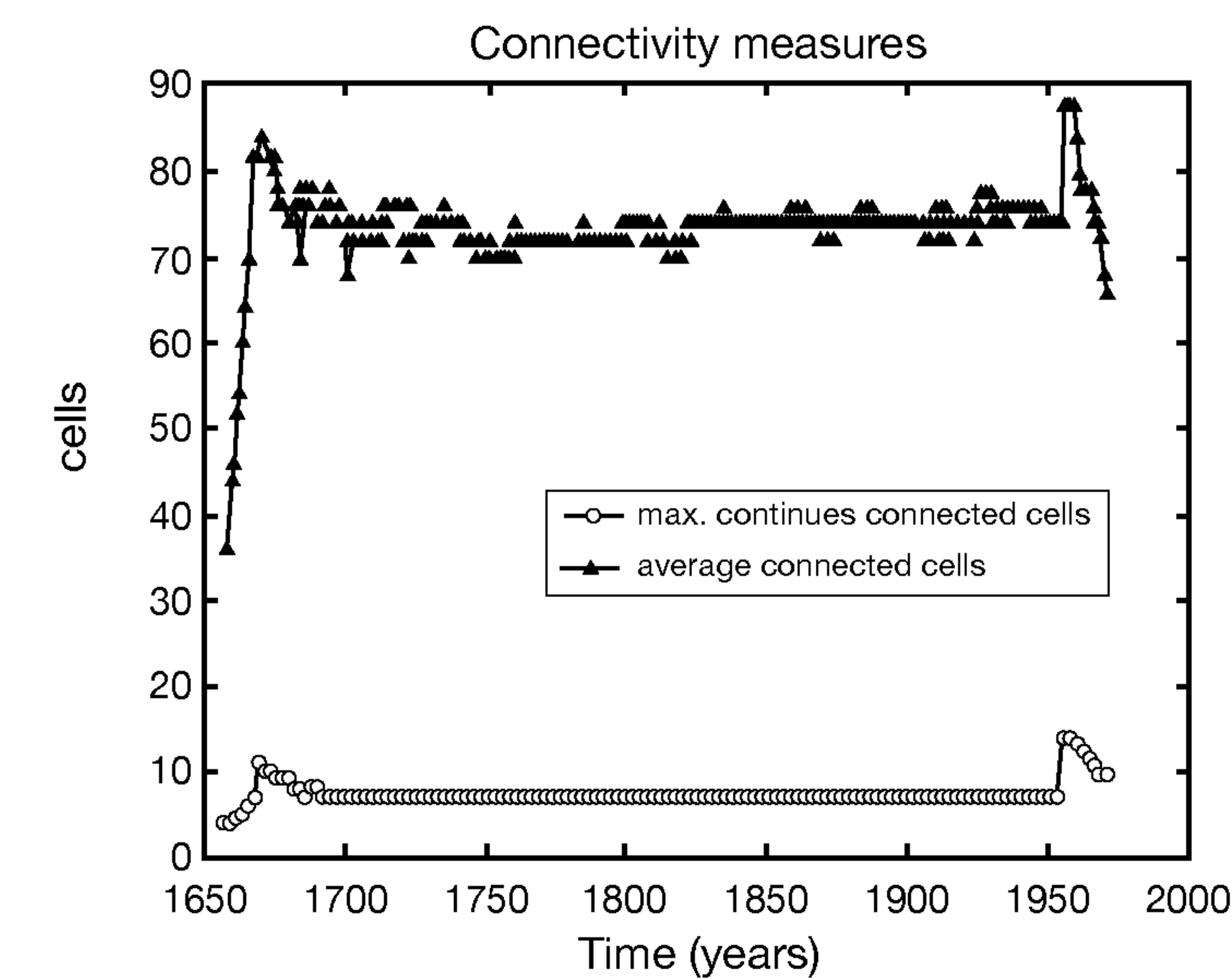
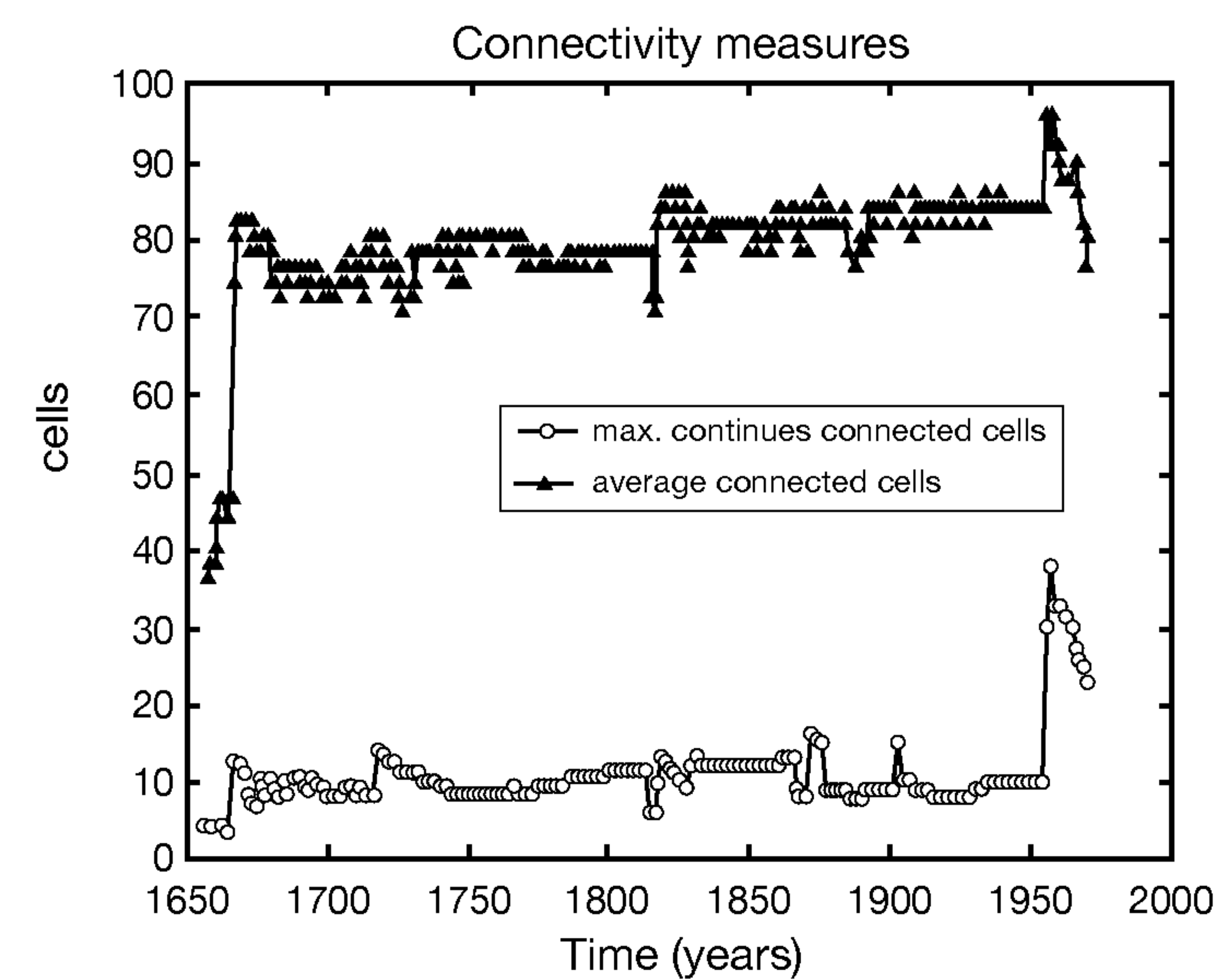
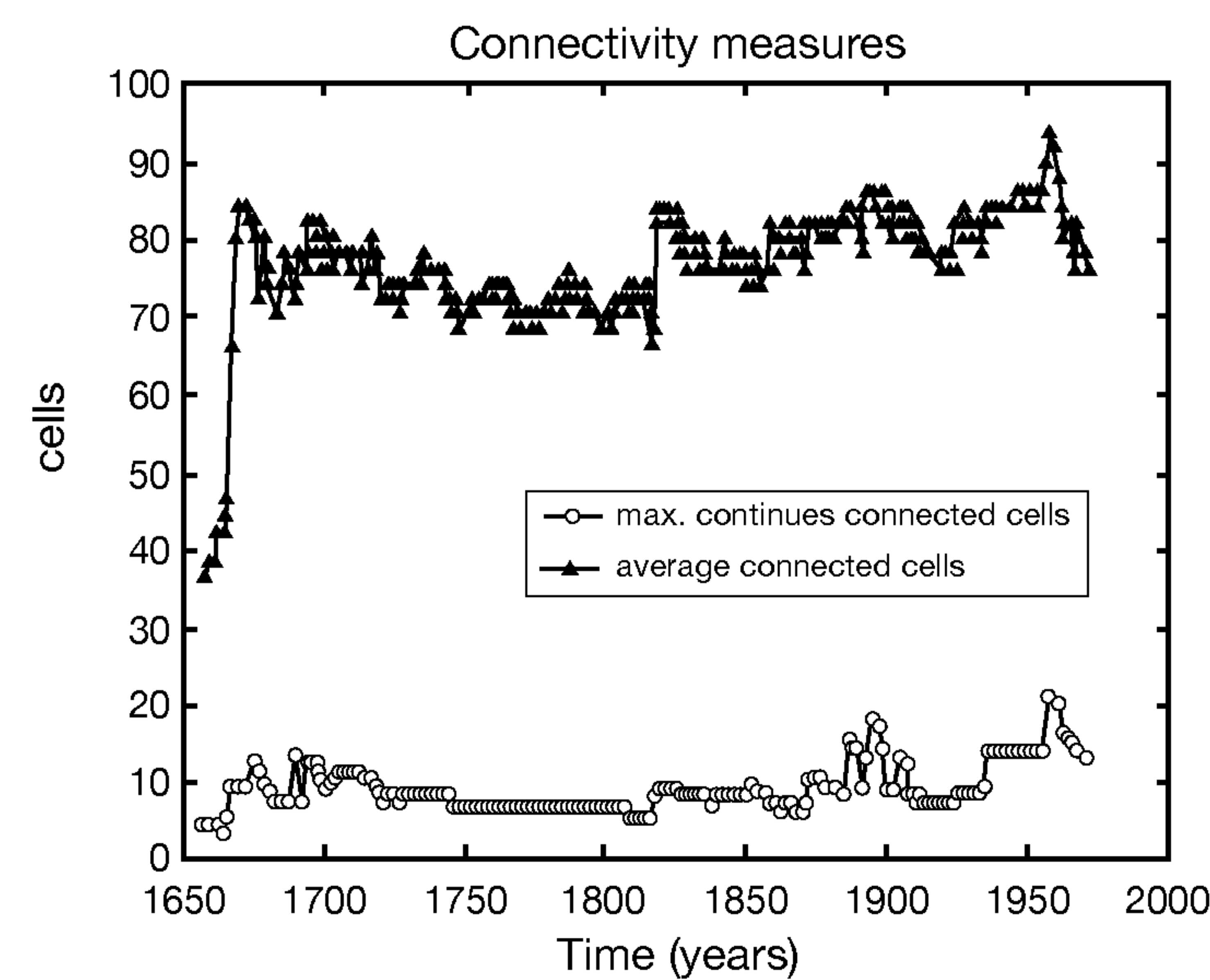
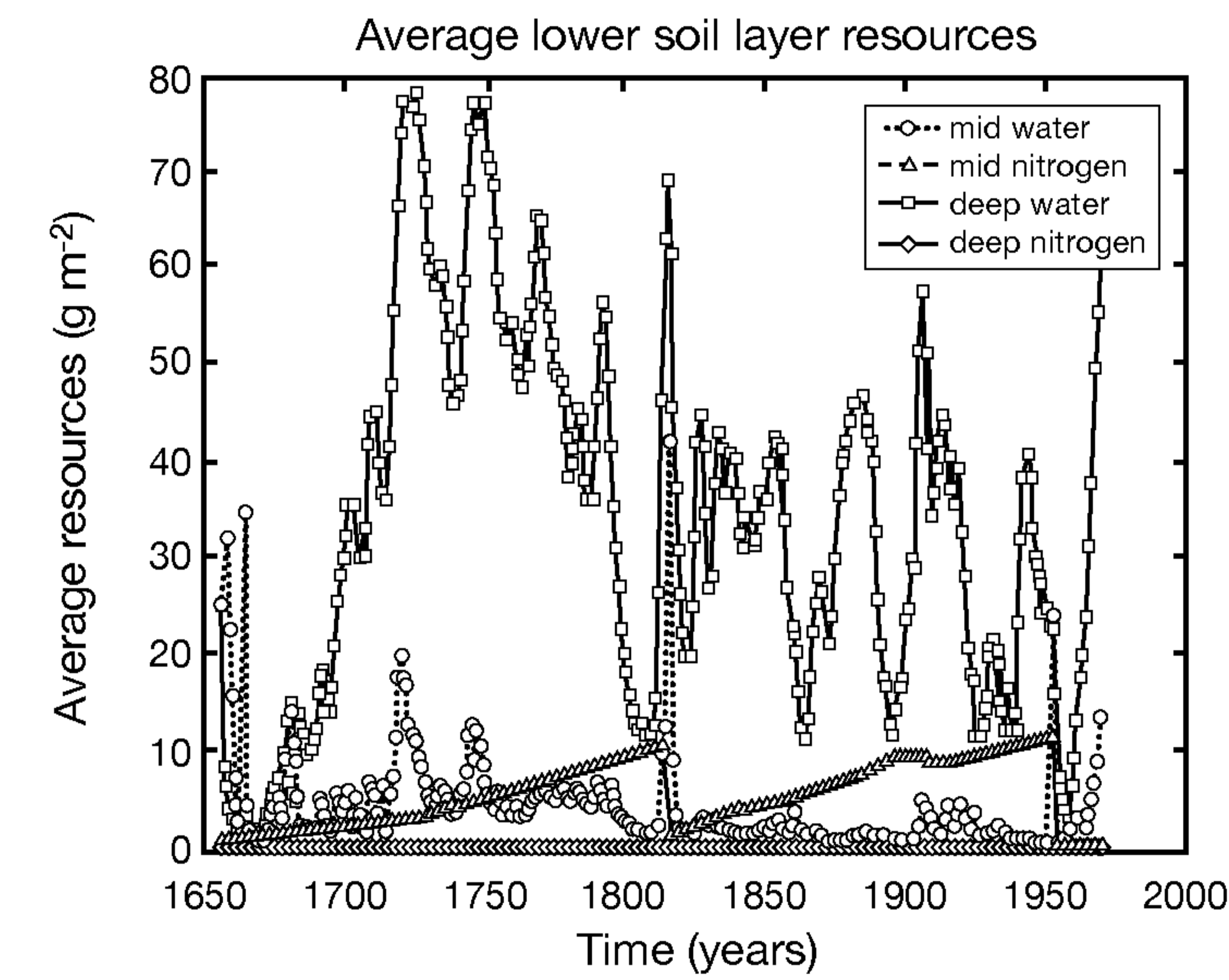
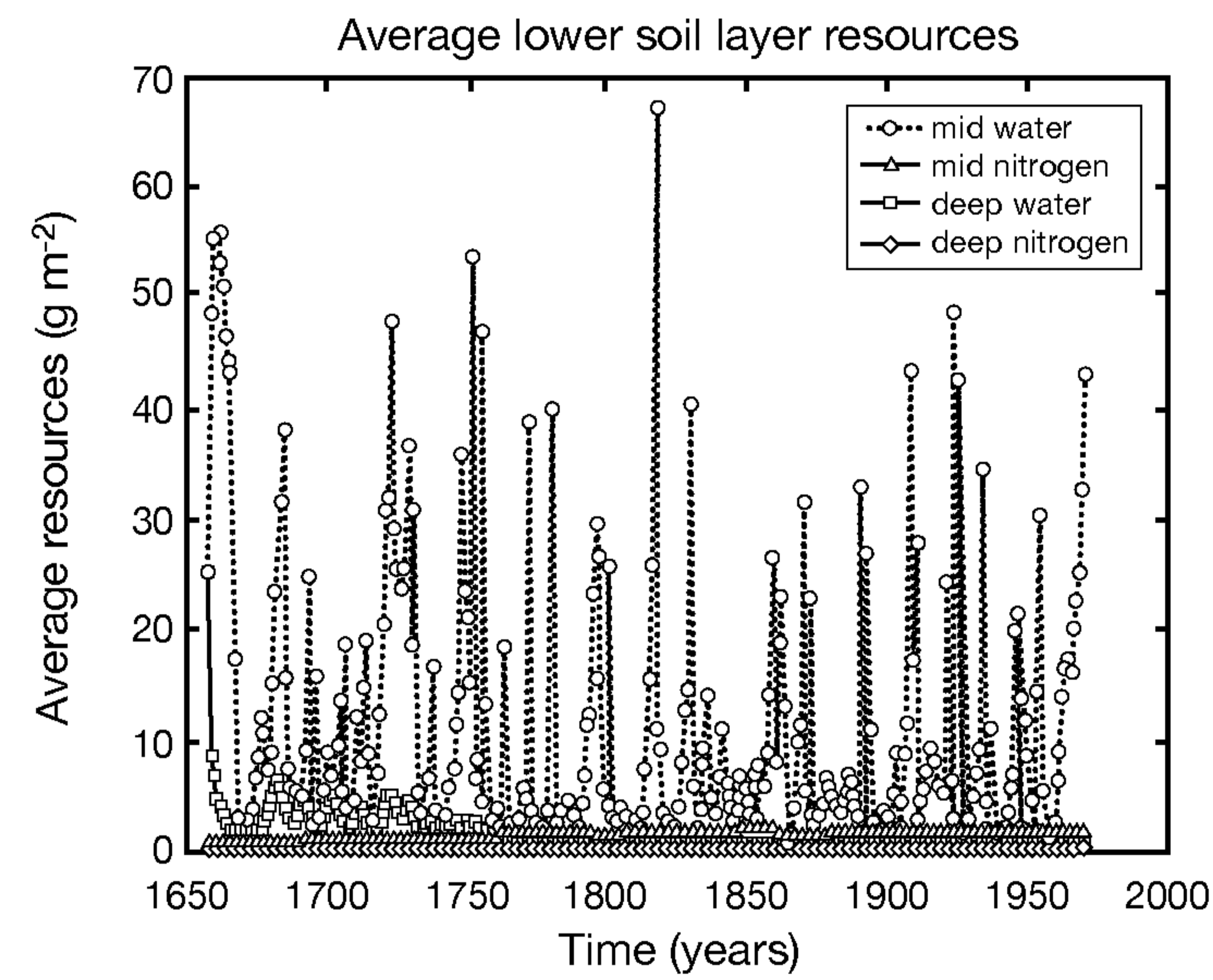
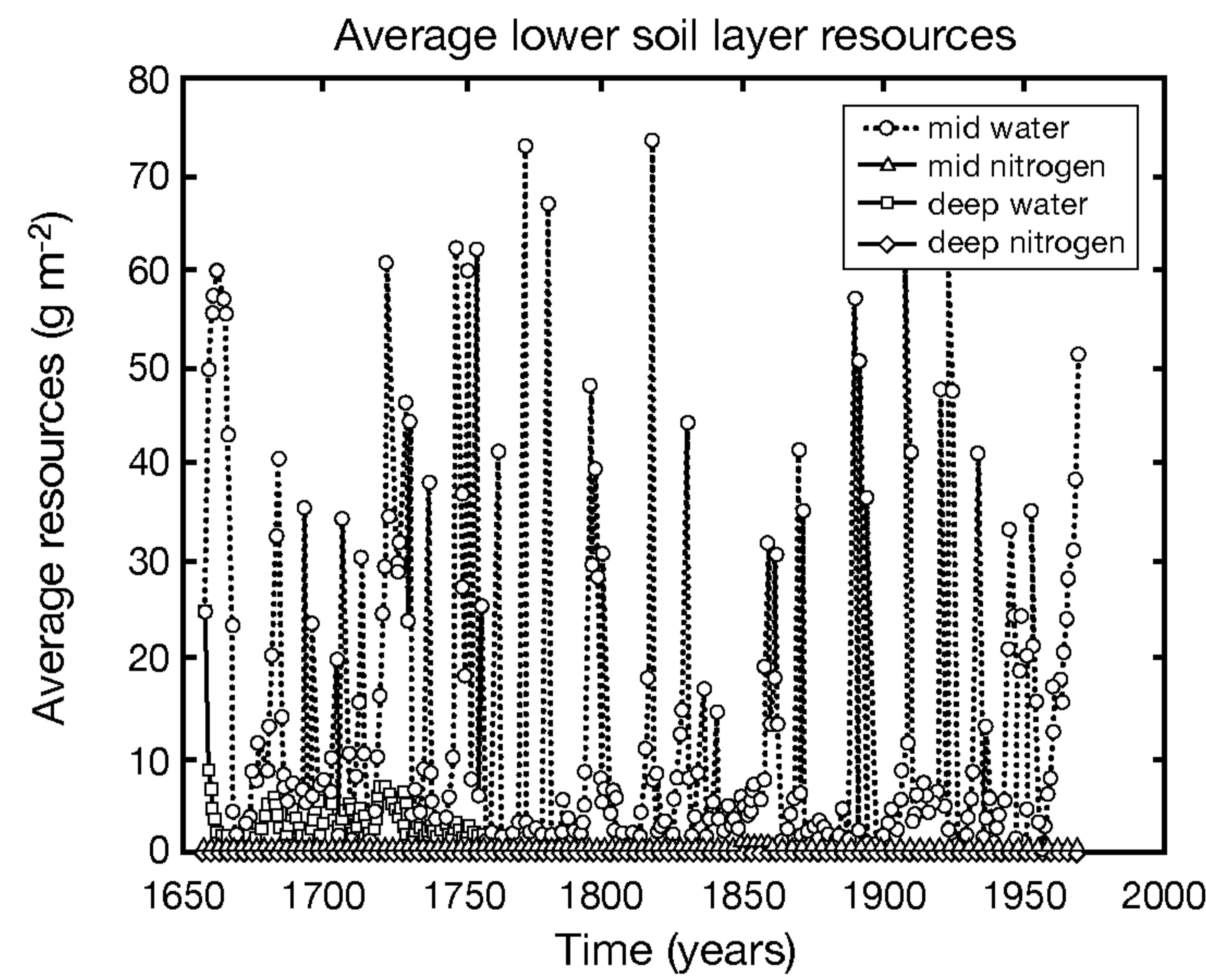
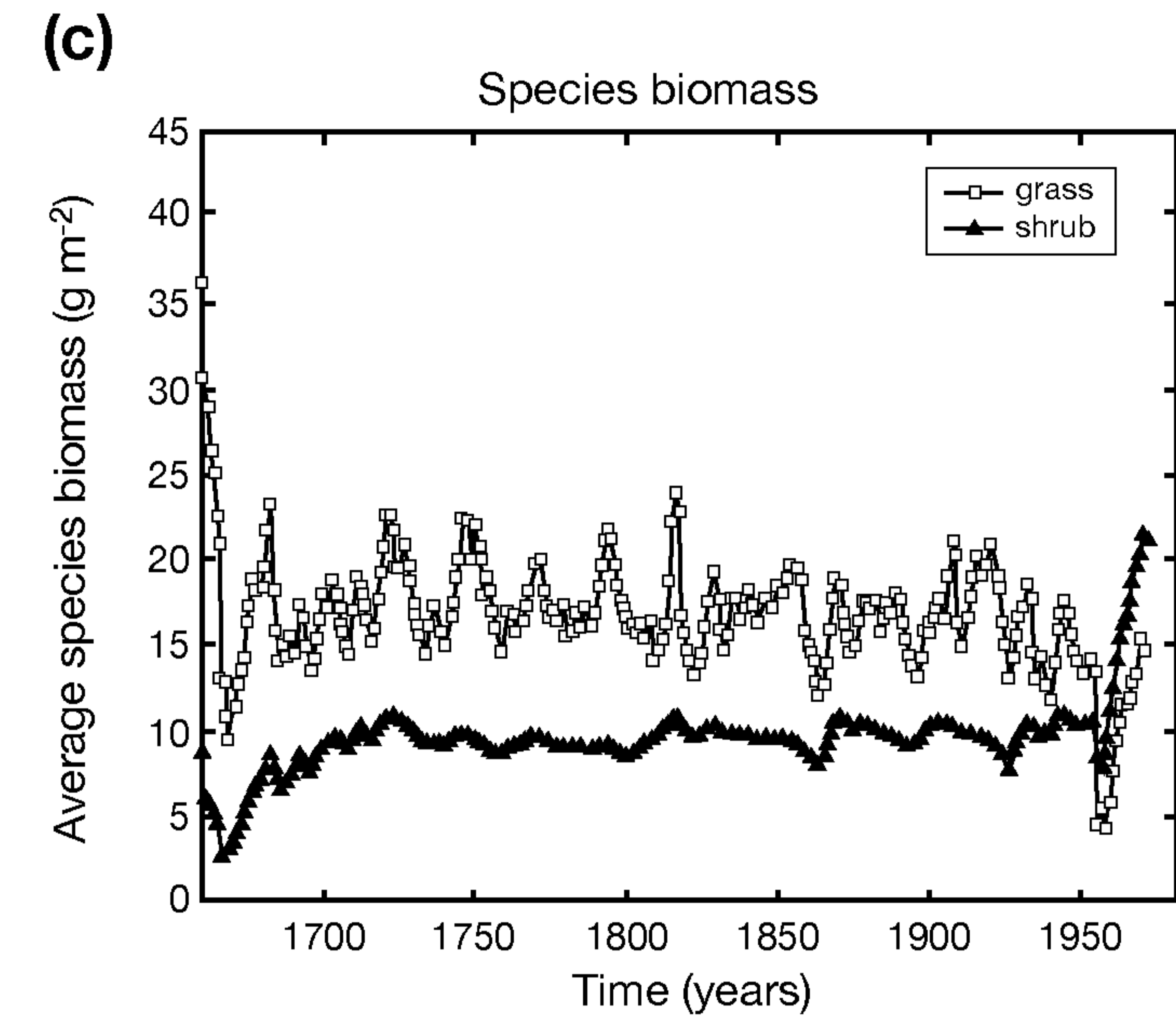
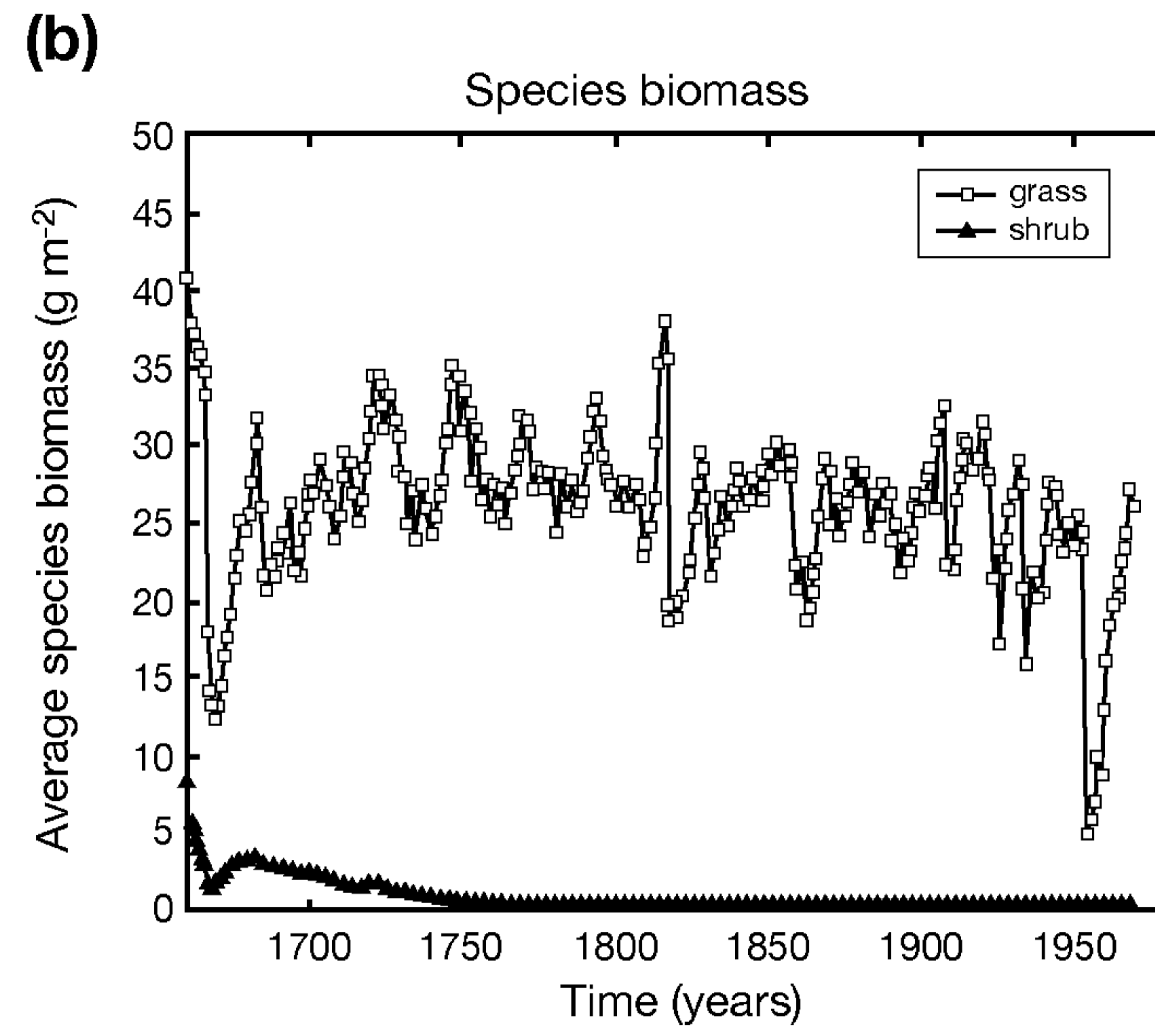
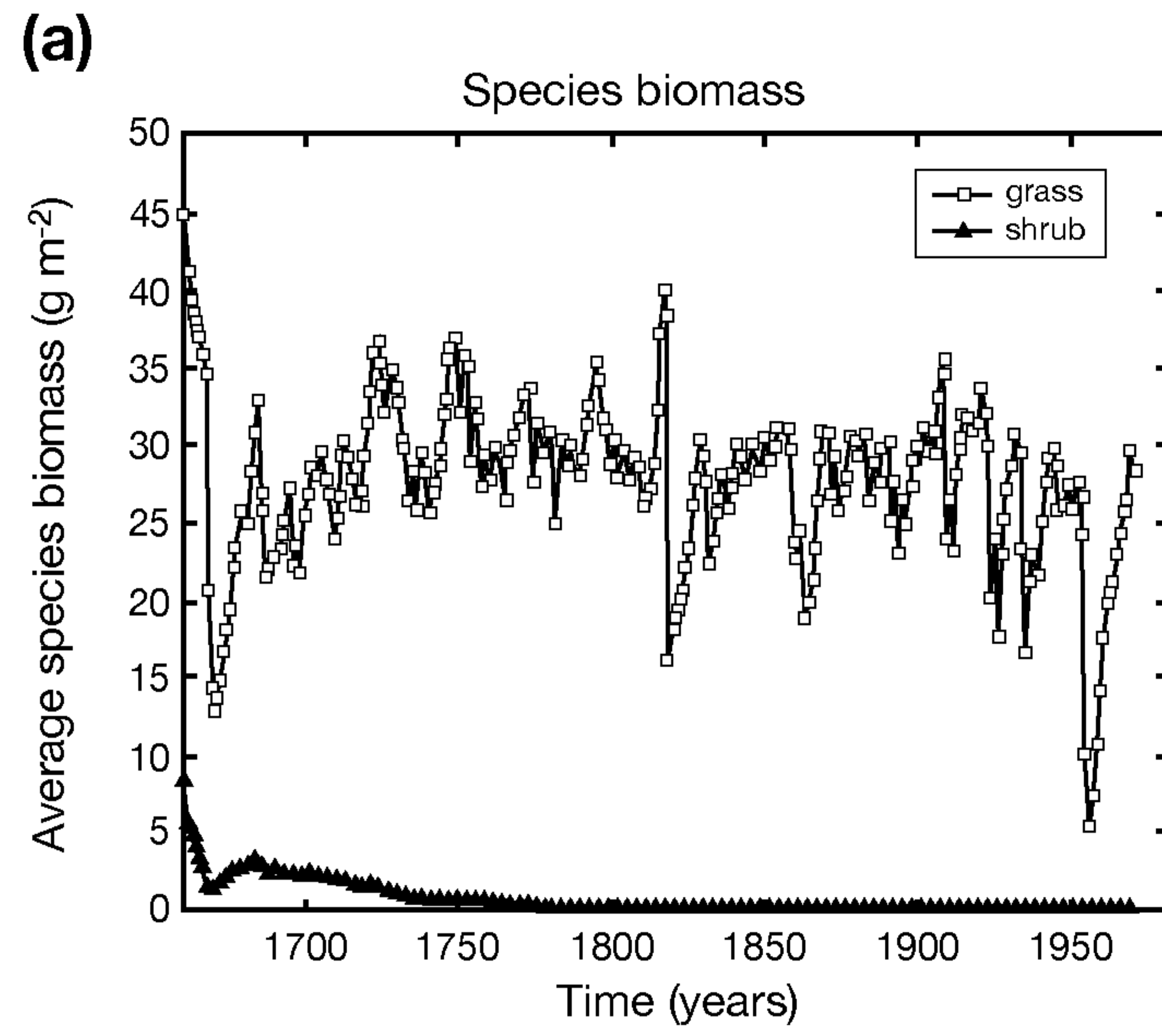


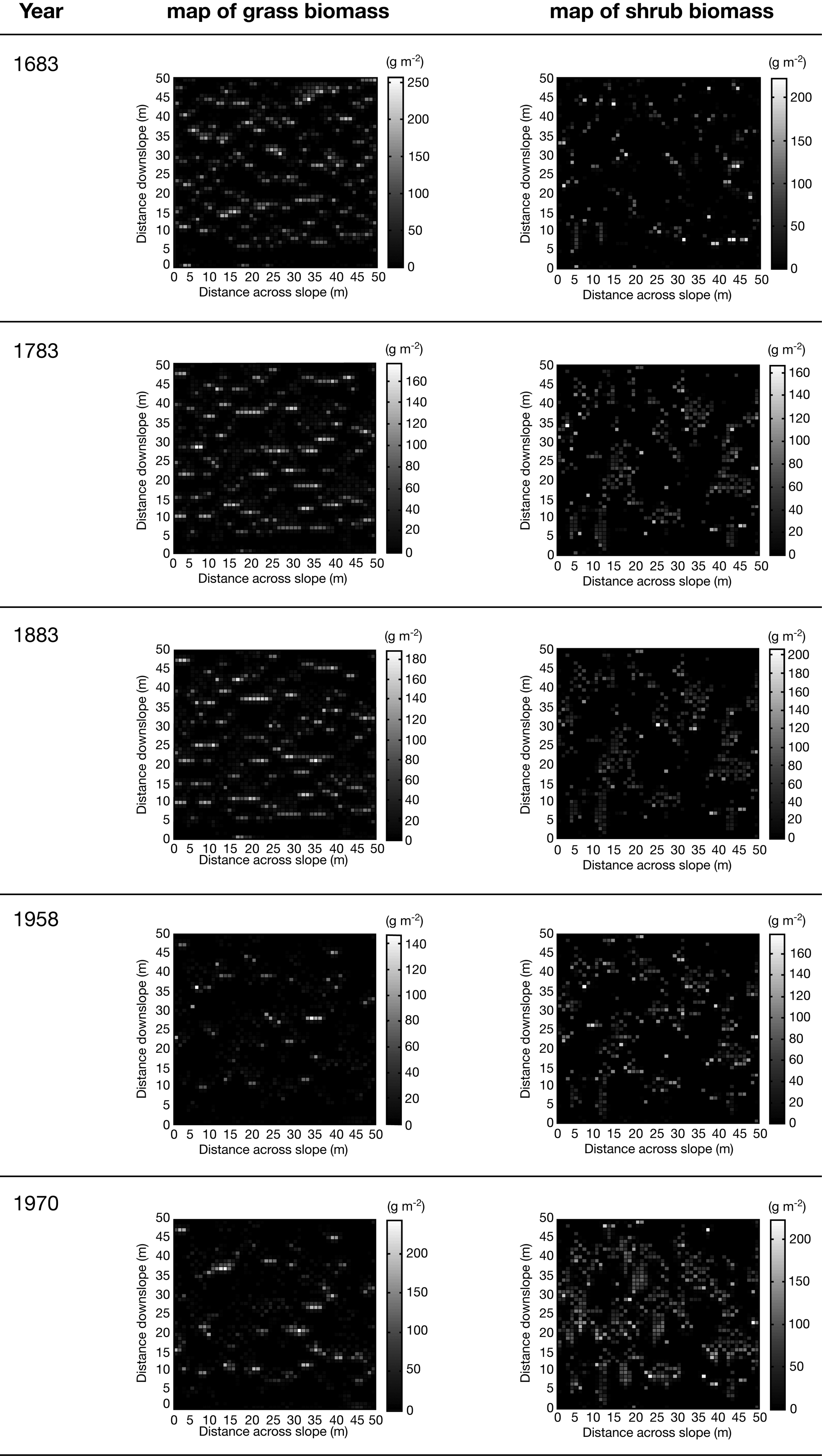
Grazing Intensity

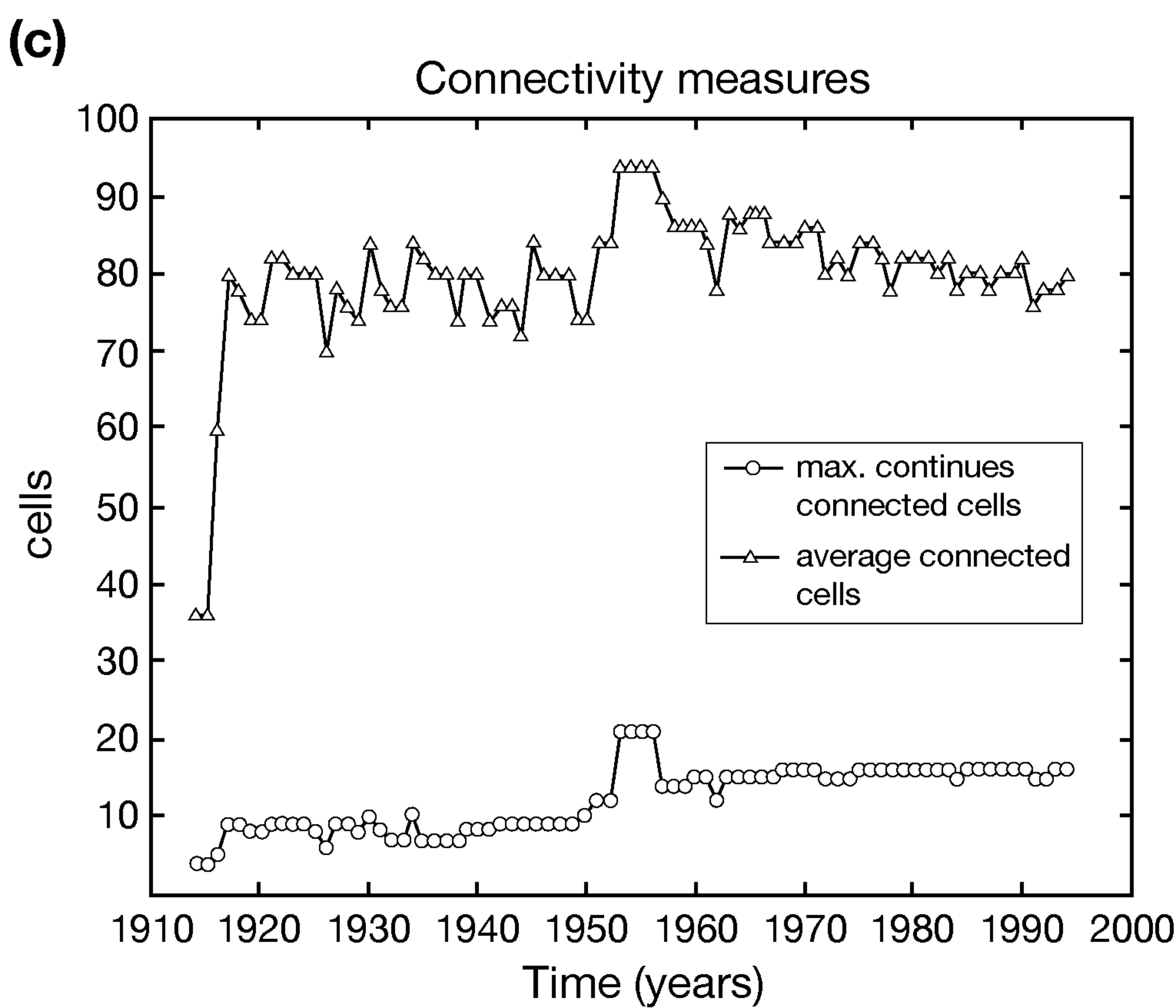
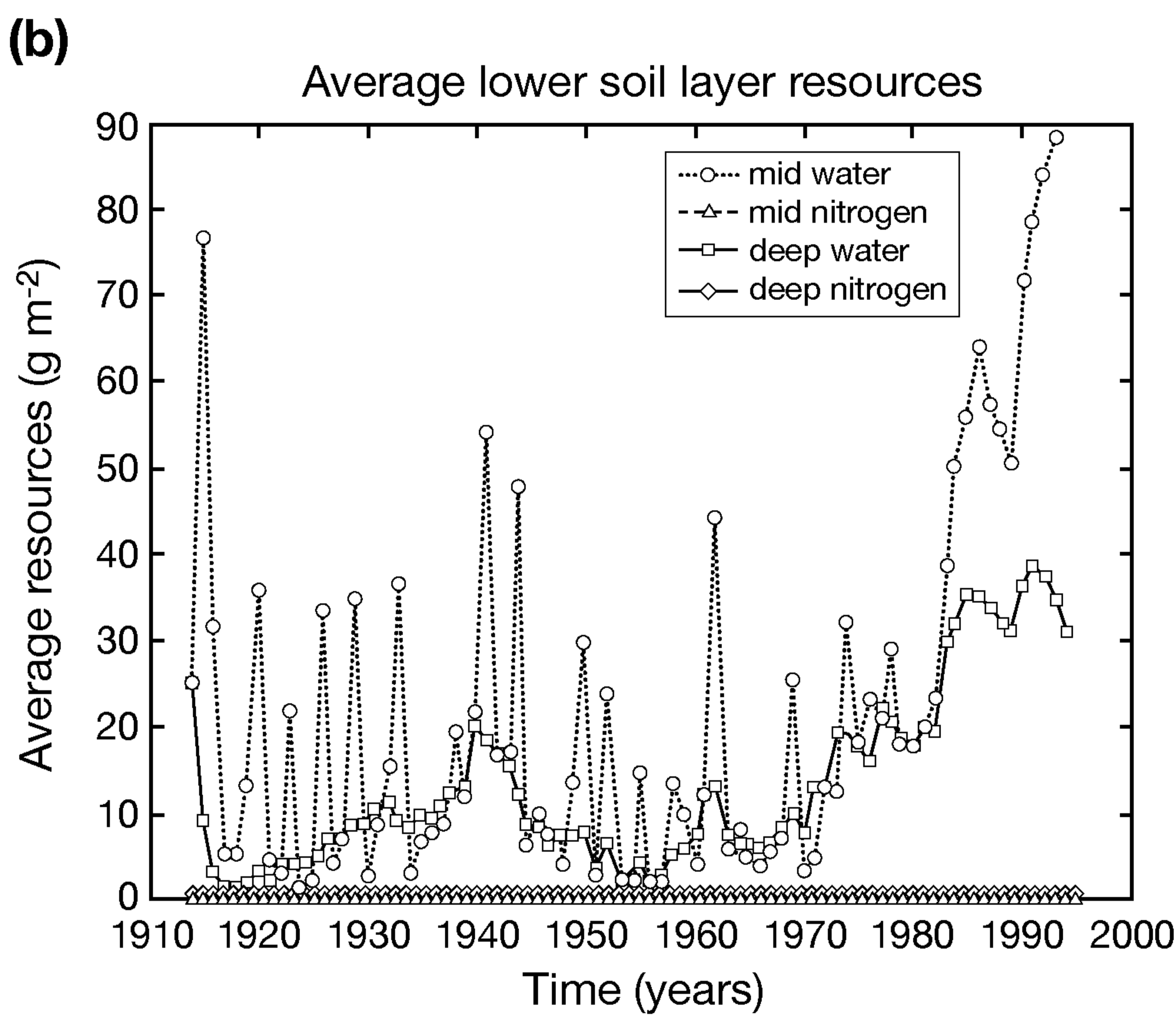
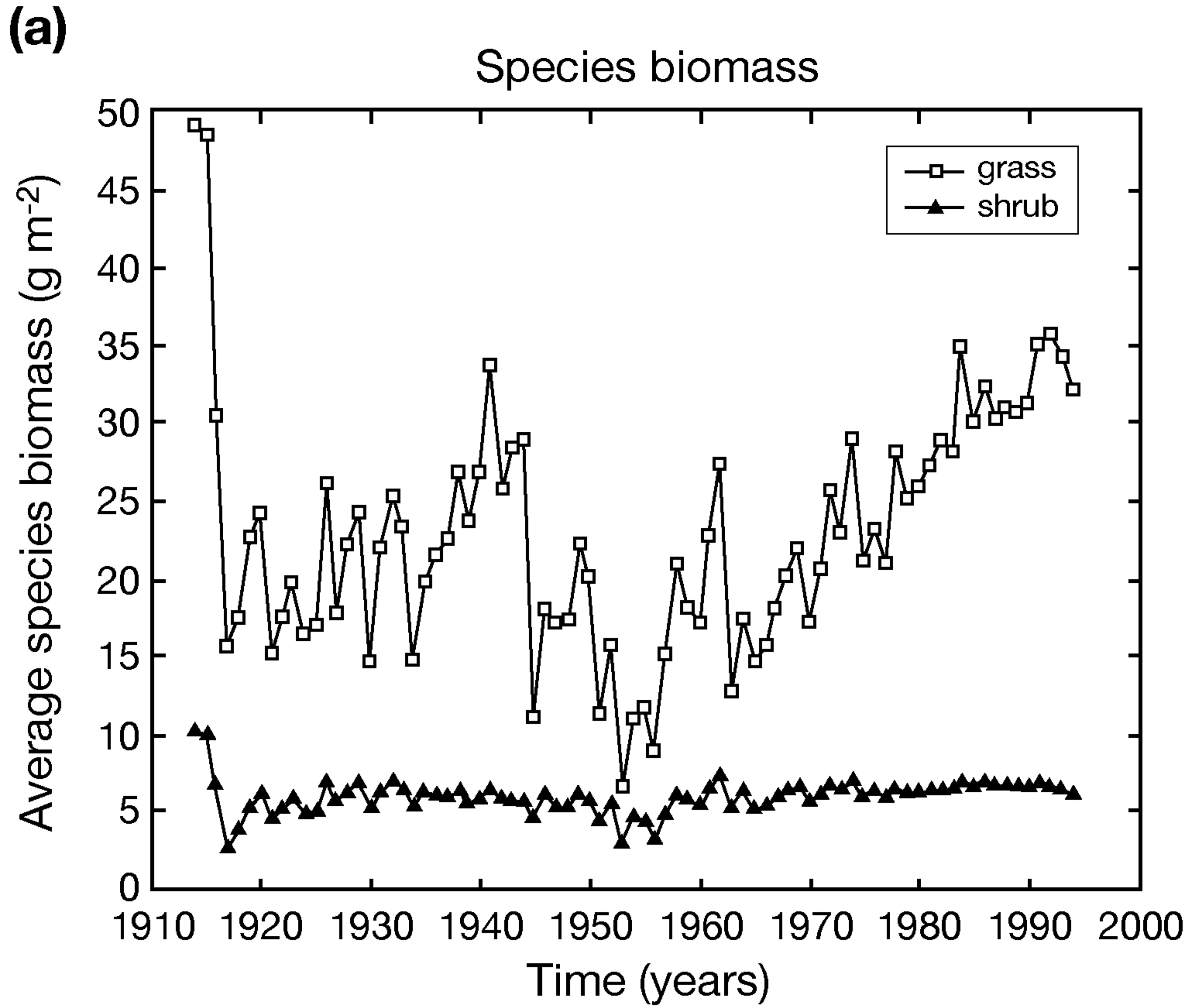
Conservative (11 g m⁻²)

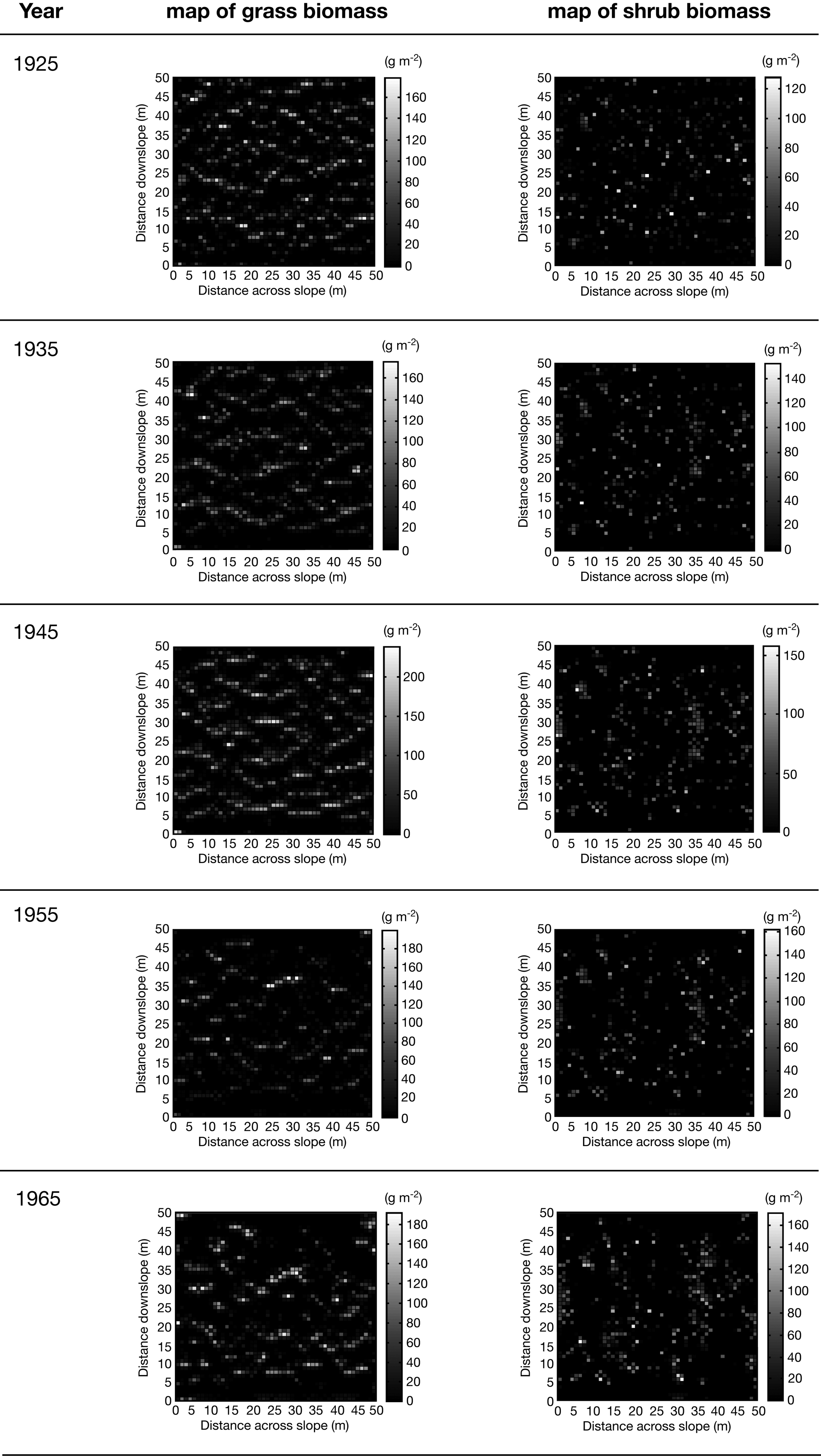
Recommended (14 g m⁻²)

Overgrazed (45 g m⁻²)







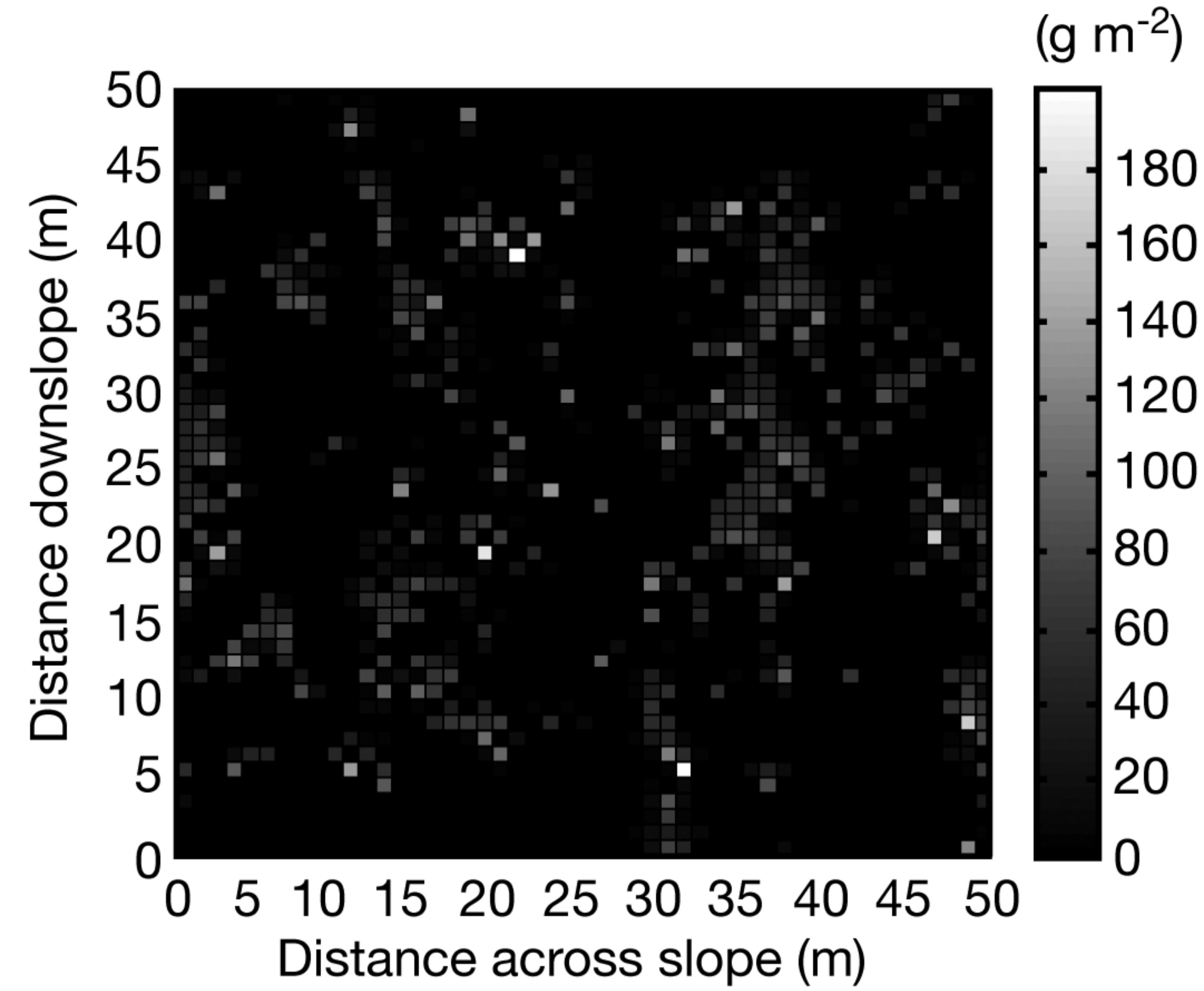
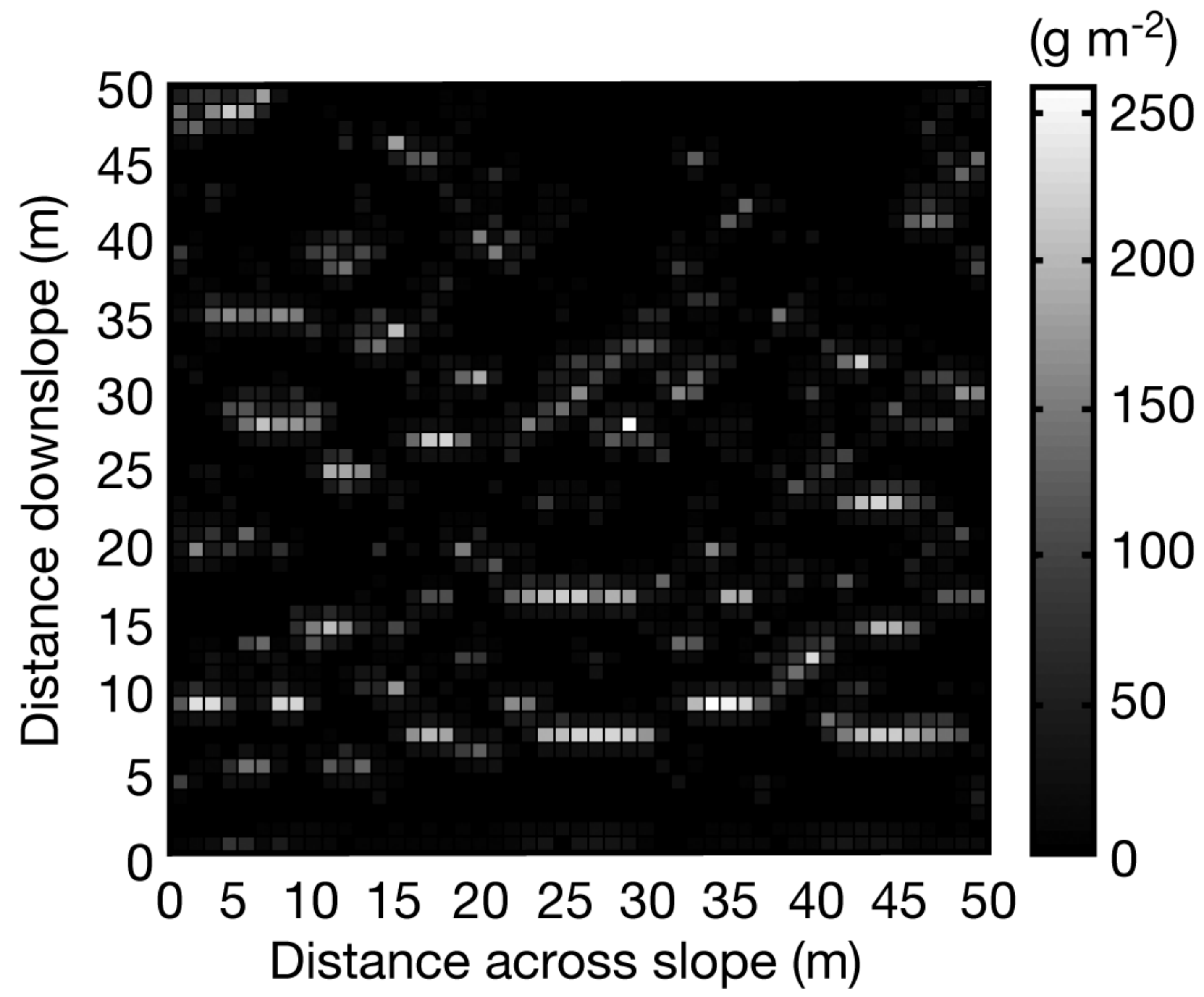


Year

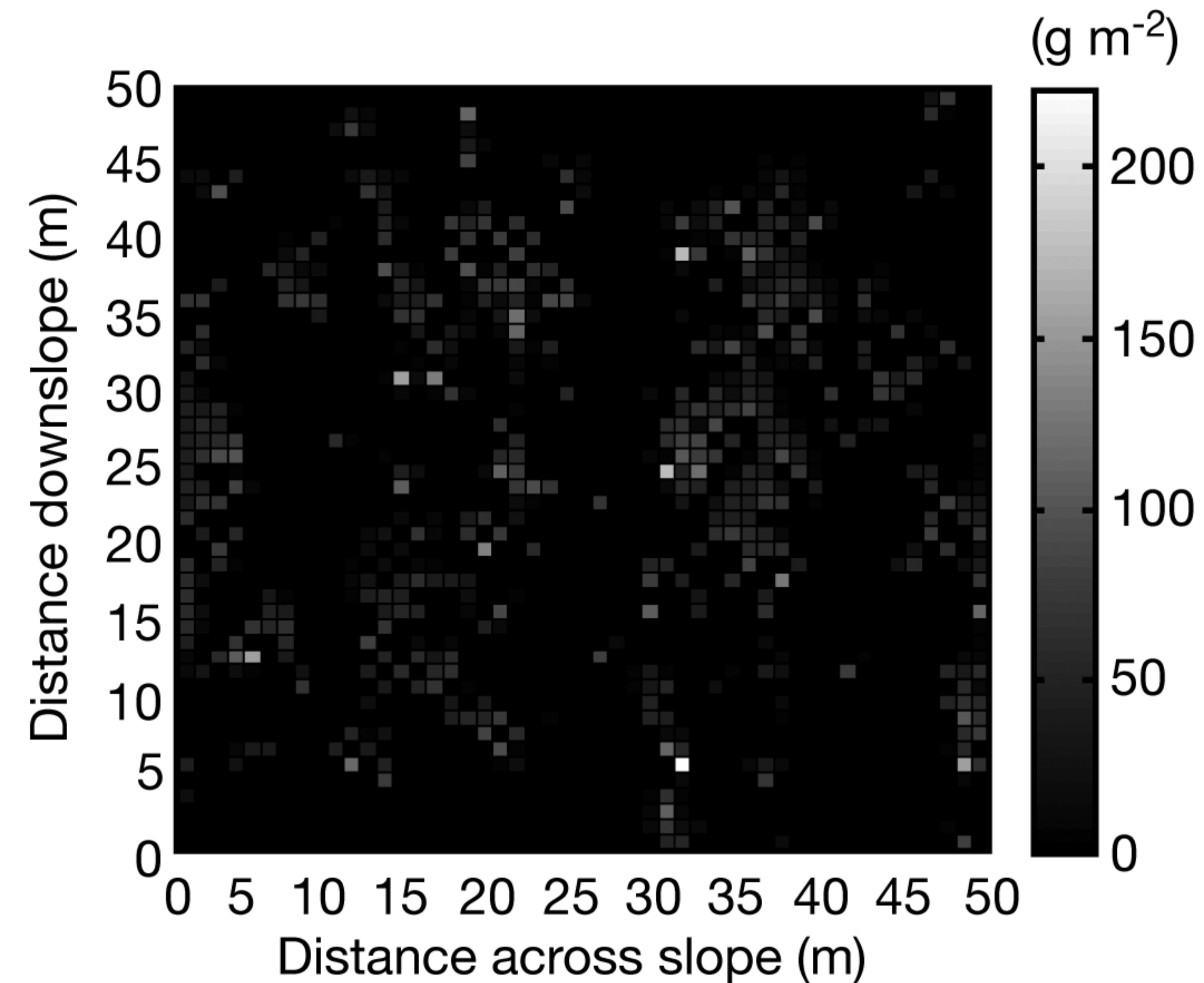
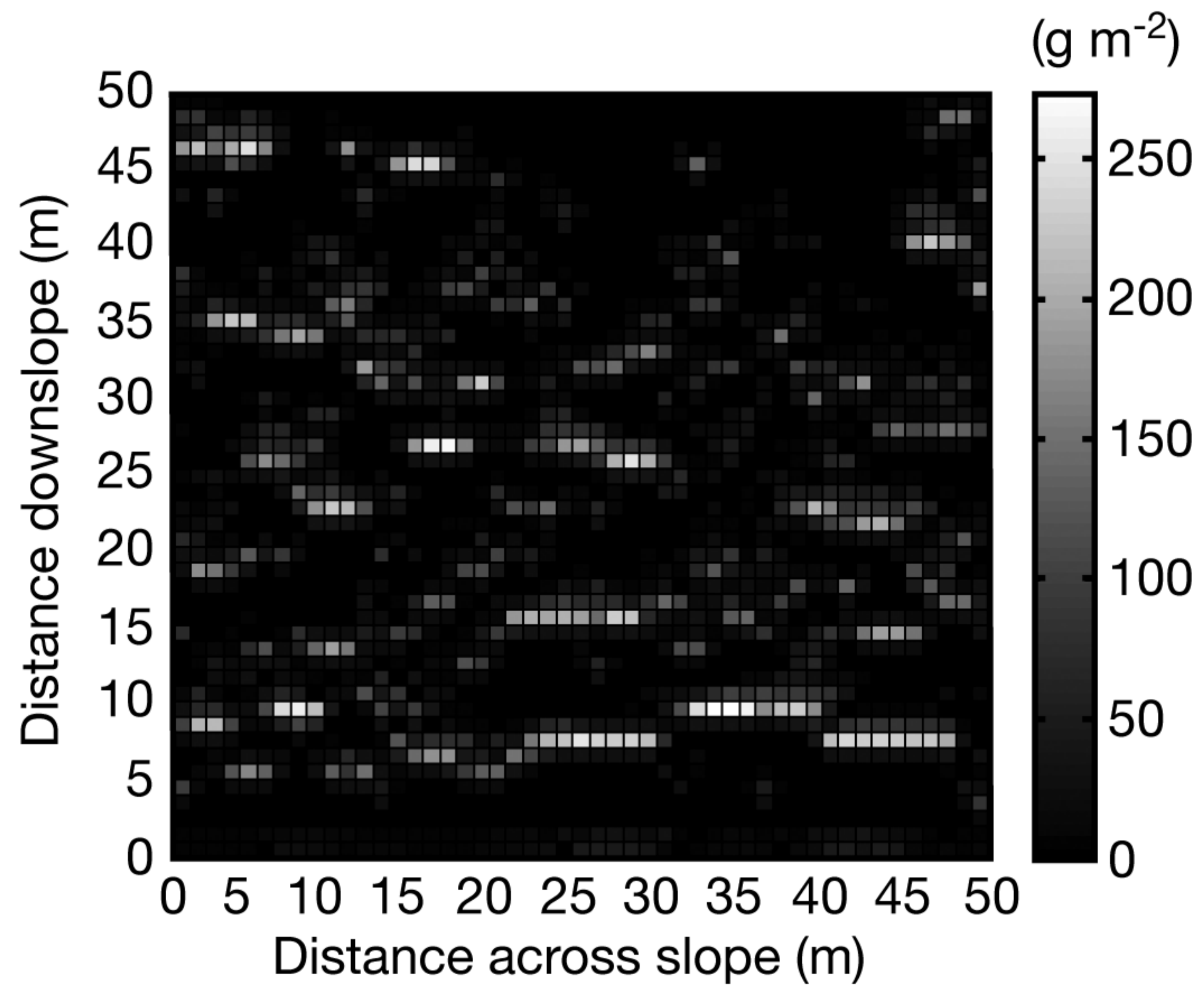
map of grass biomass

map of shrub biomass

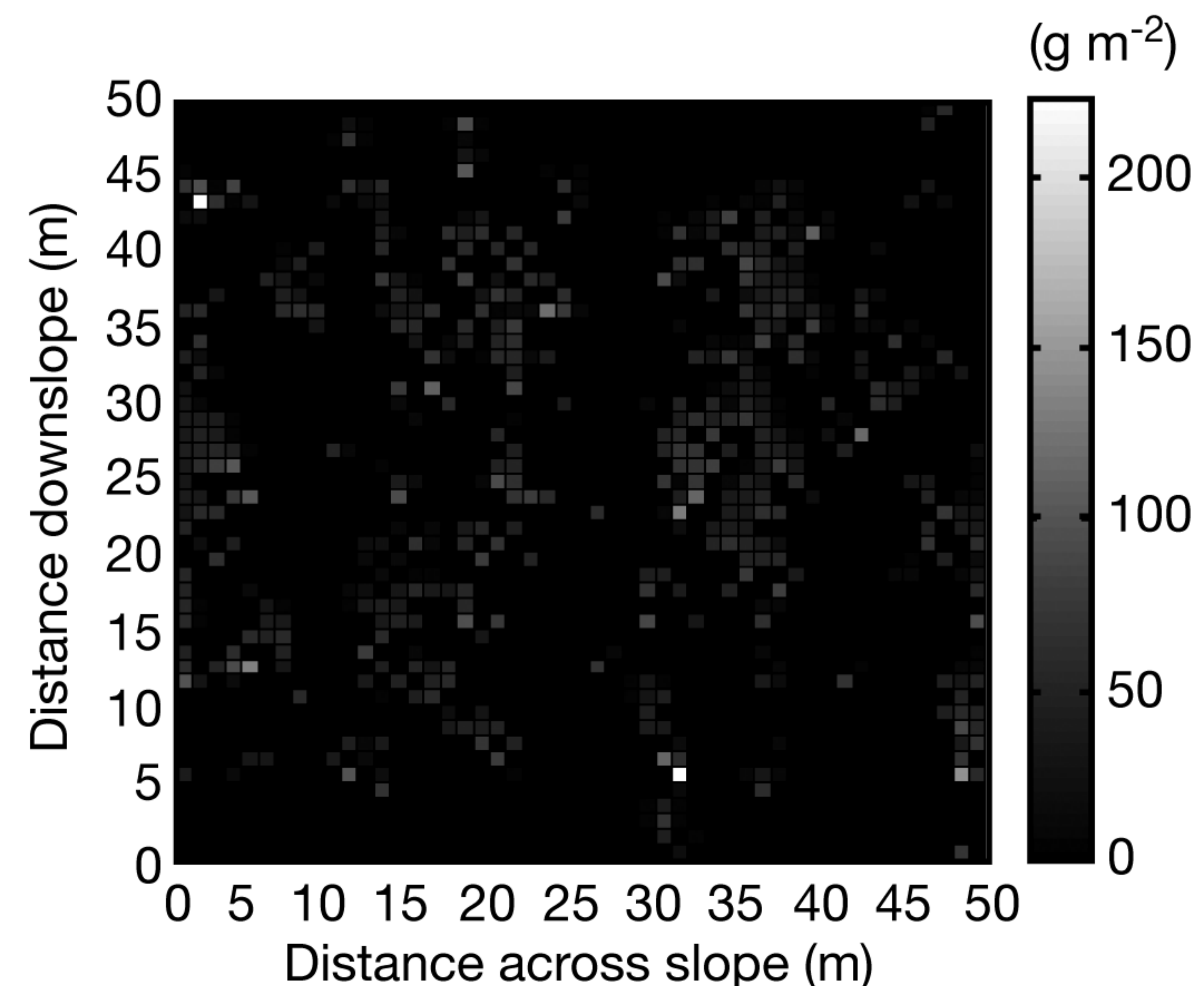
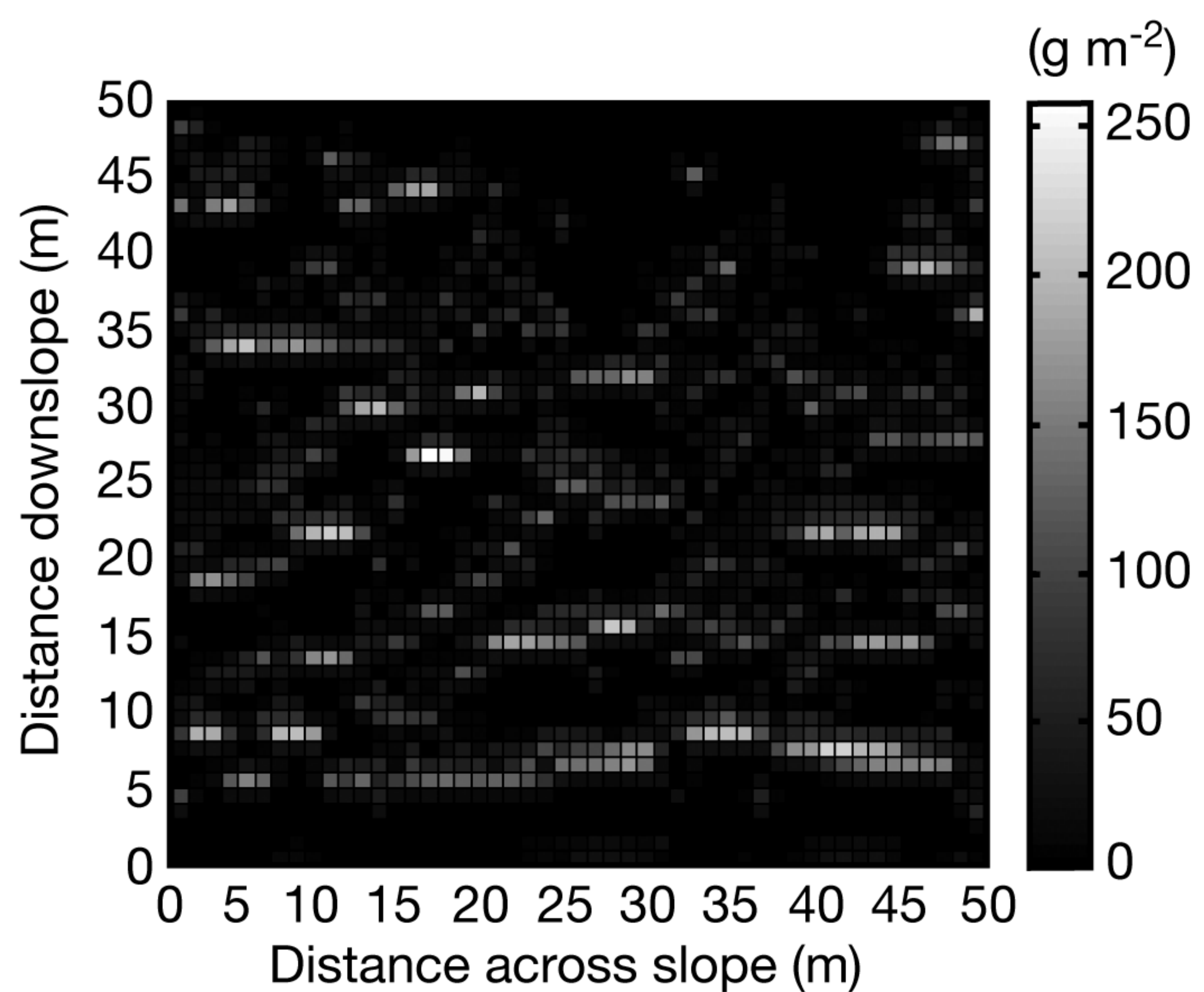
1975



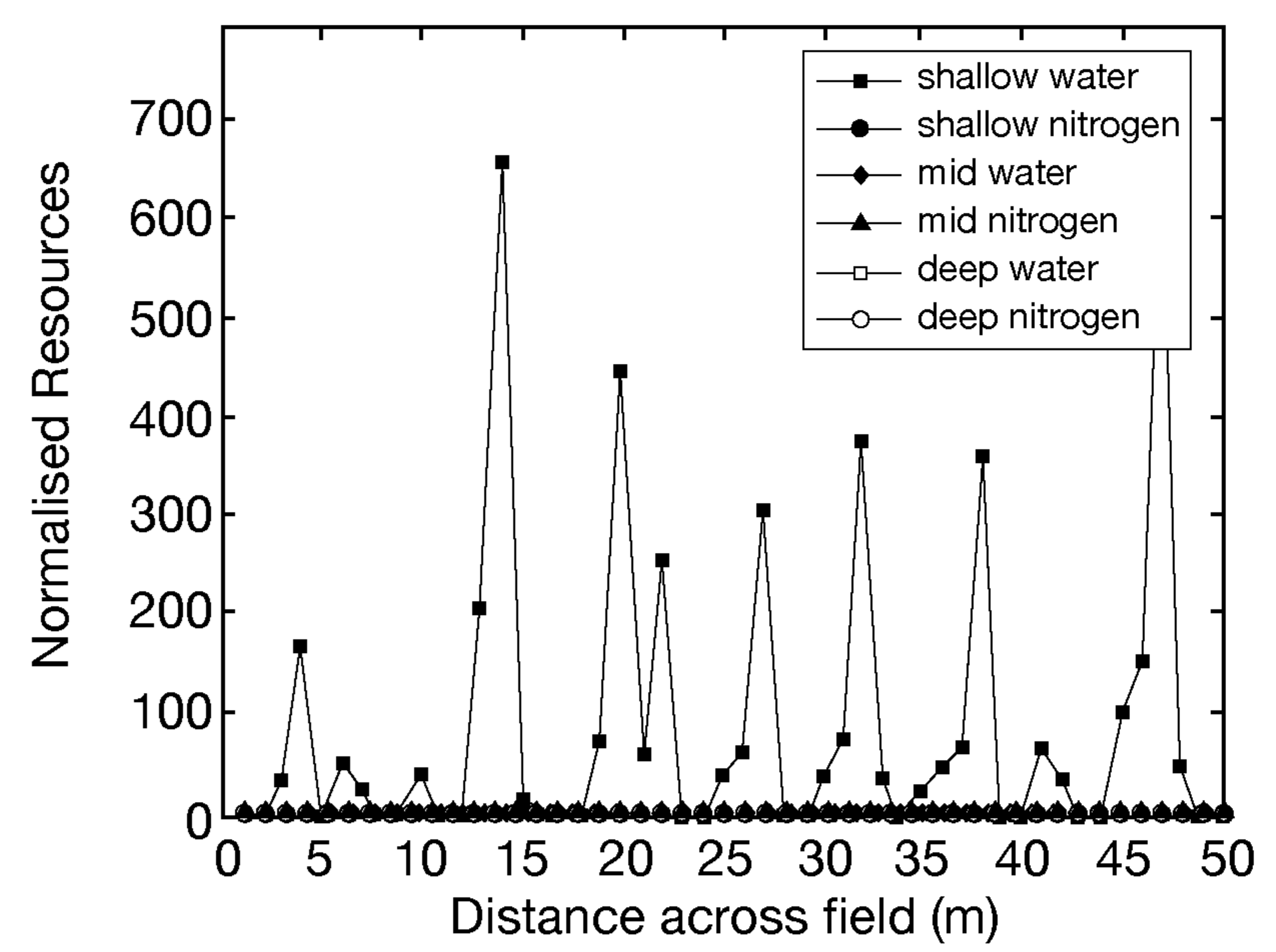
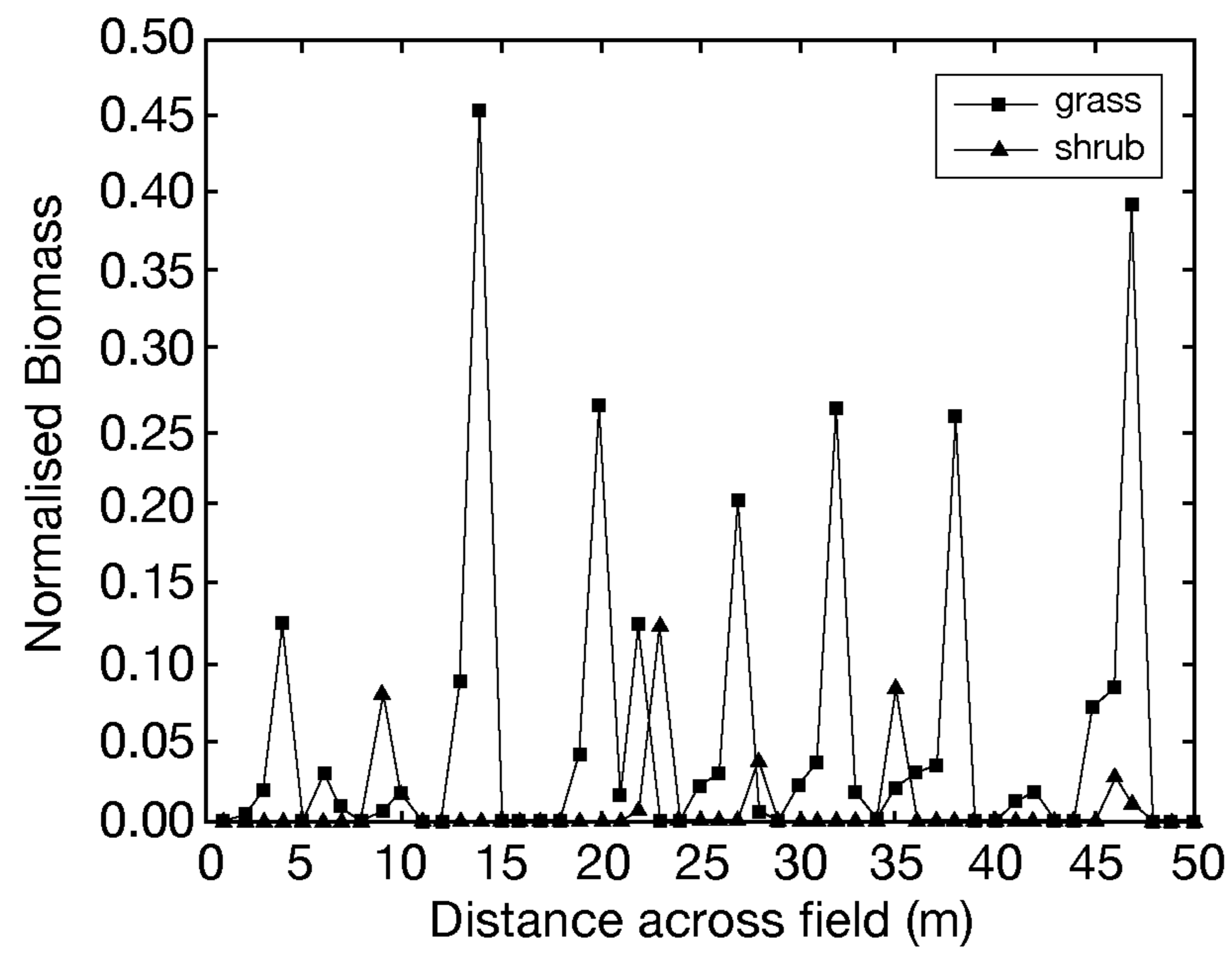
1985



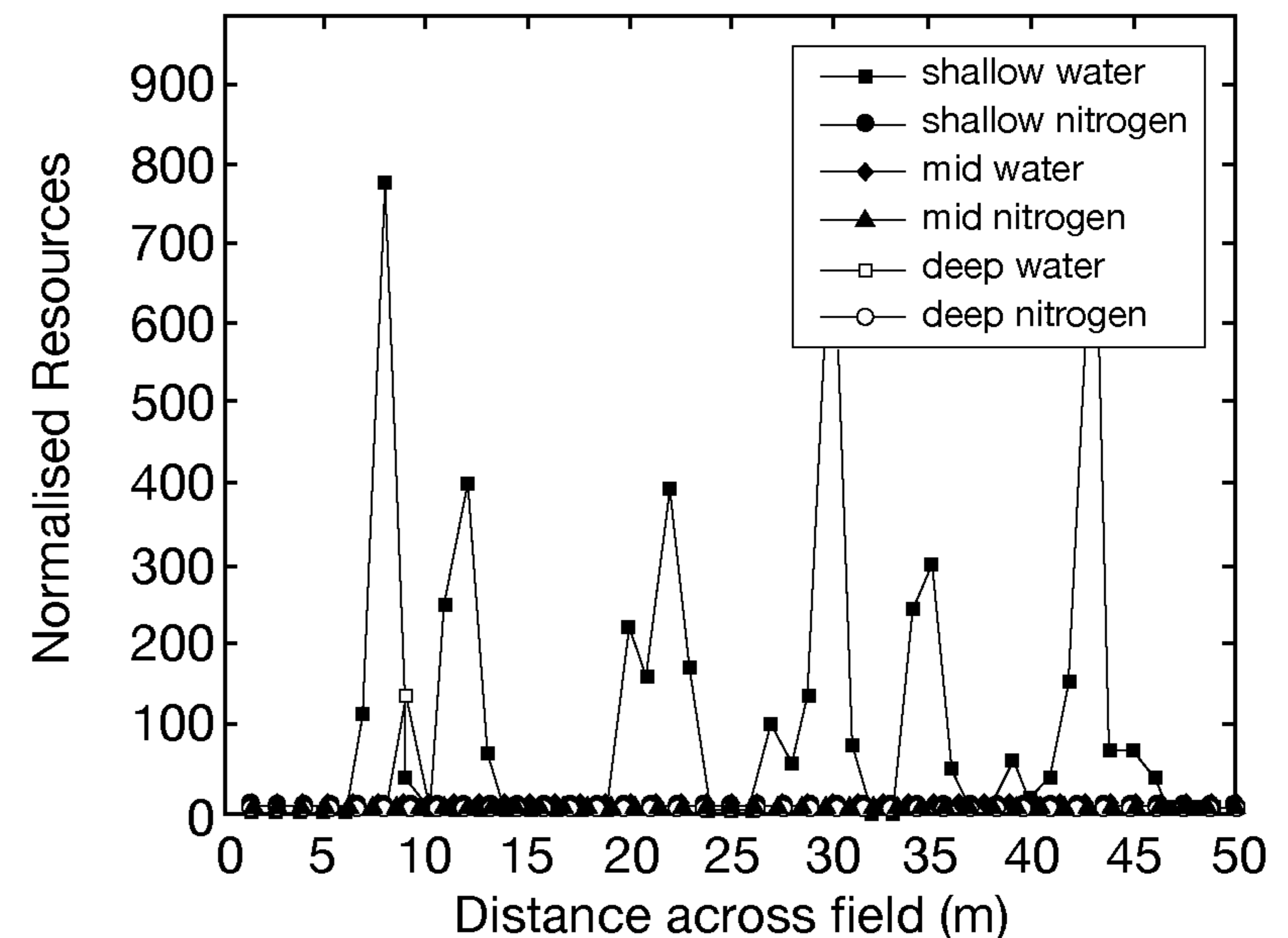
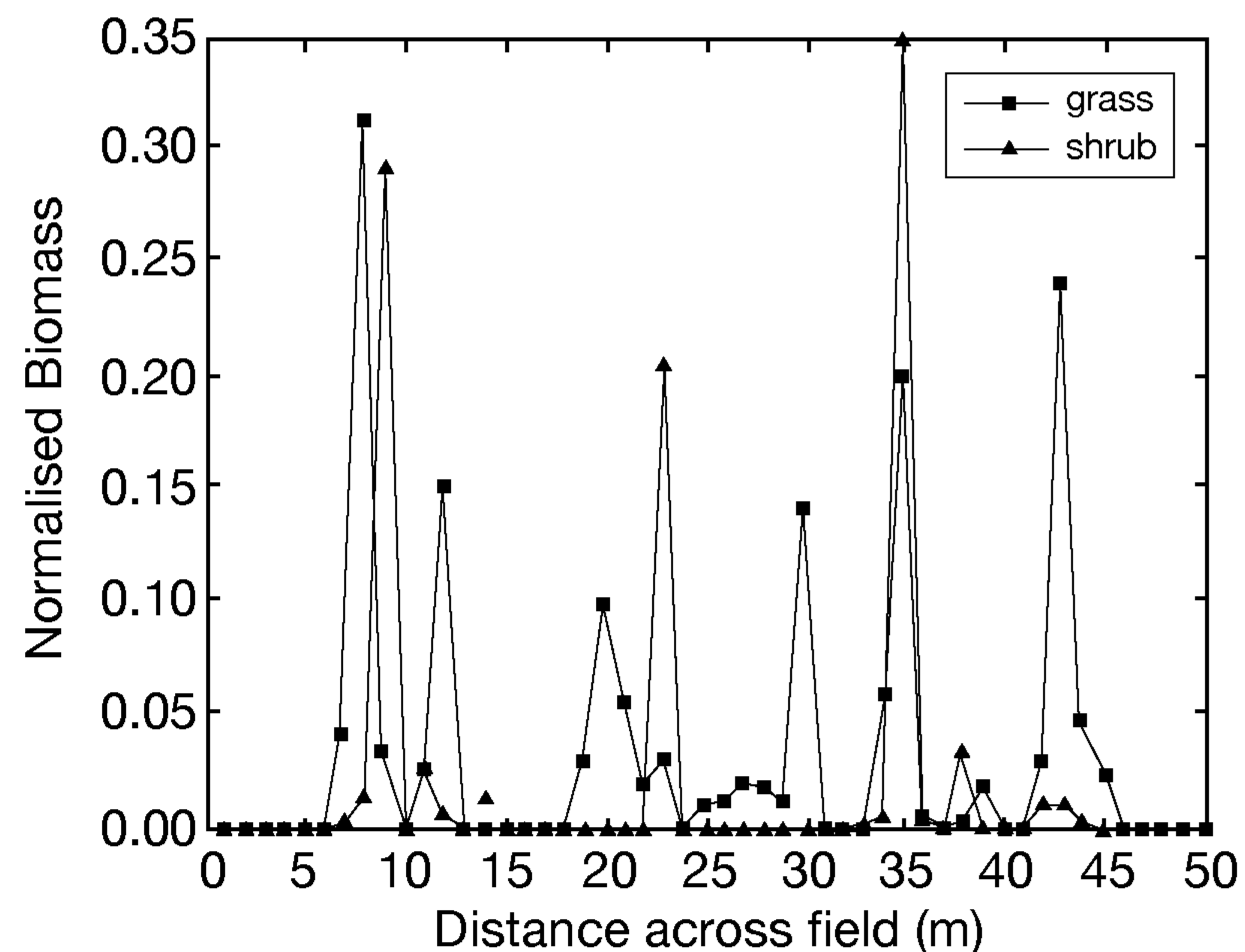
1995



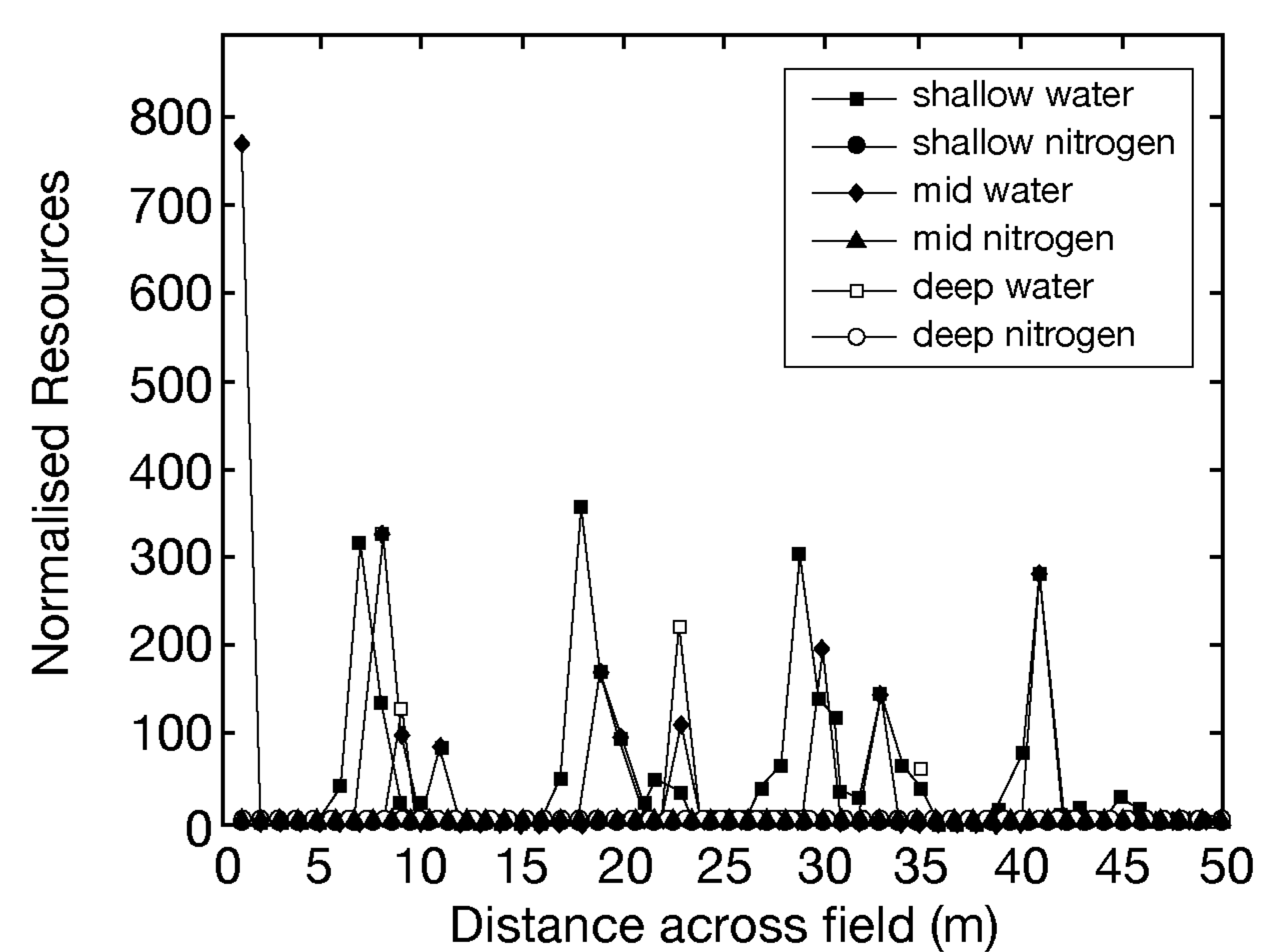
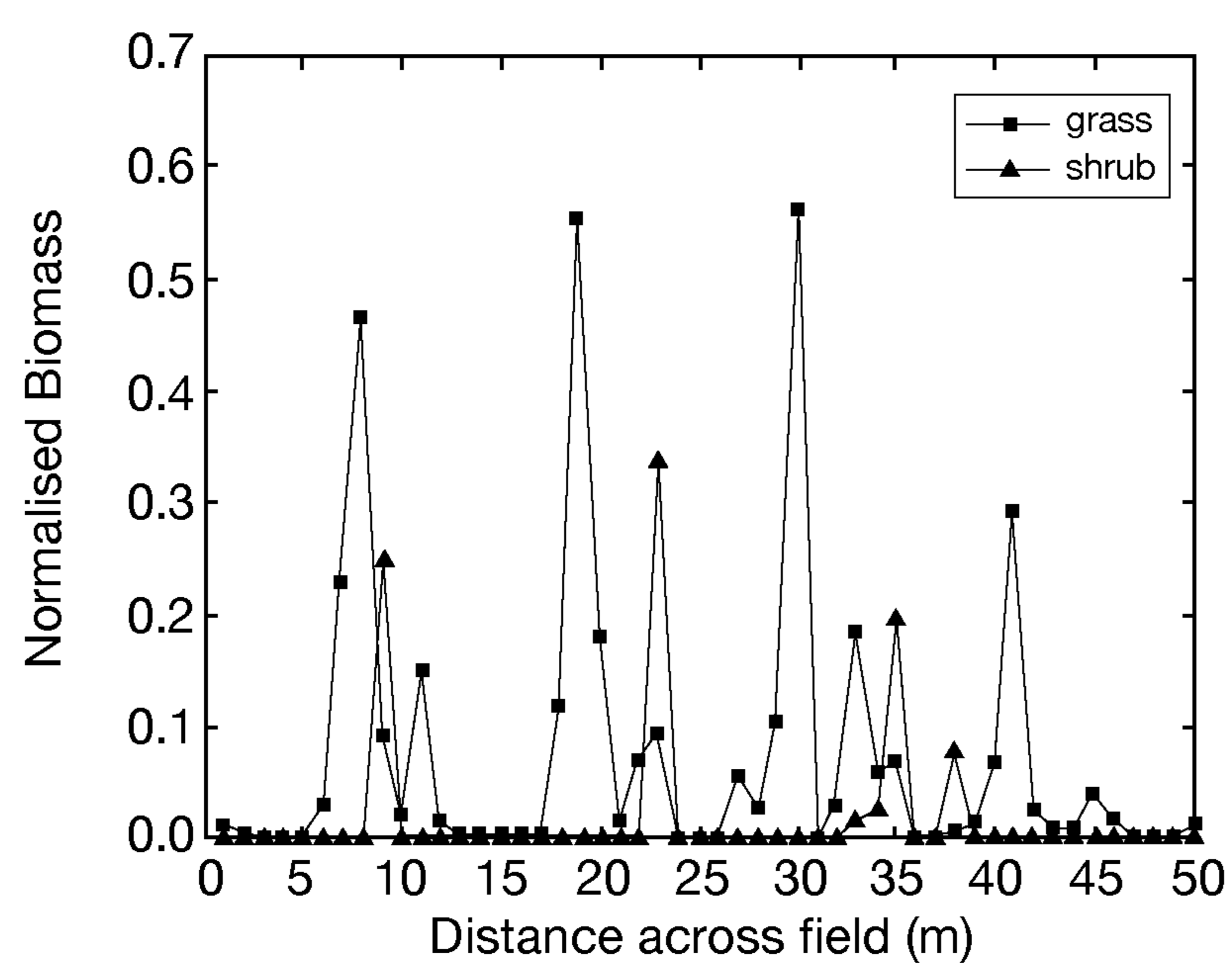
1925



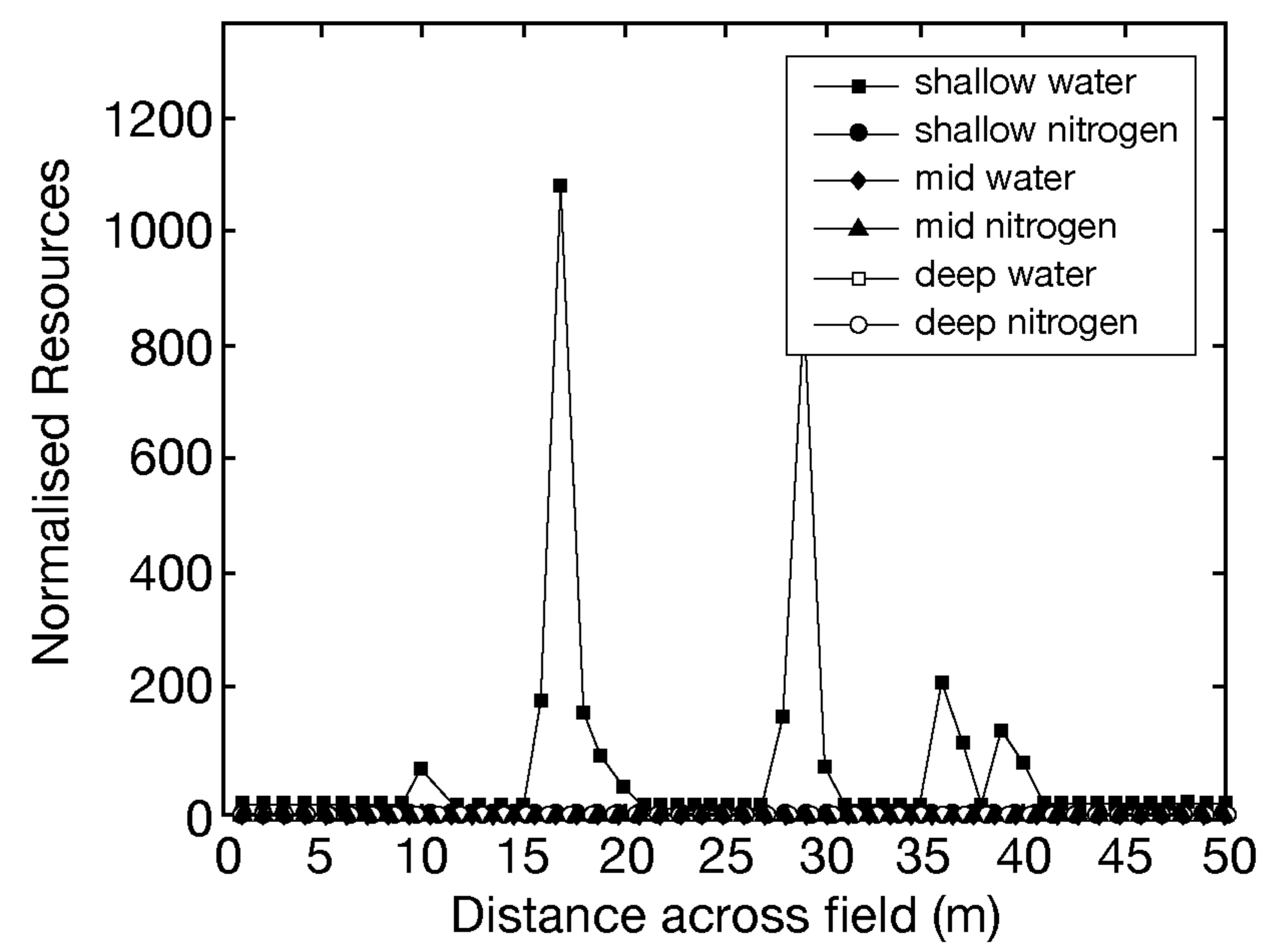
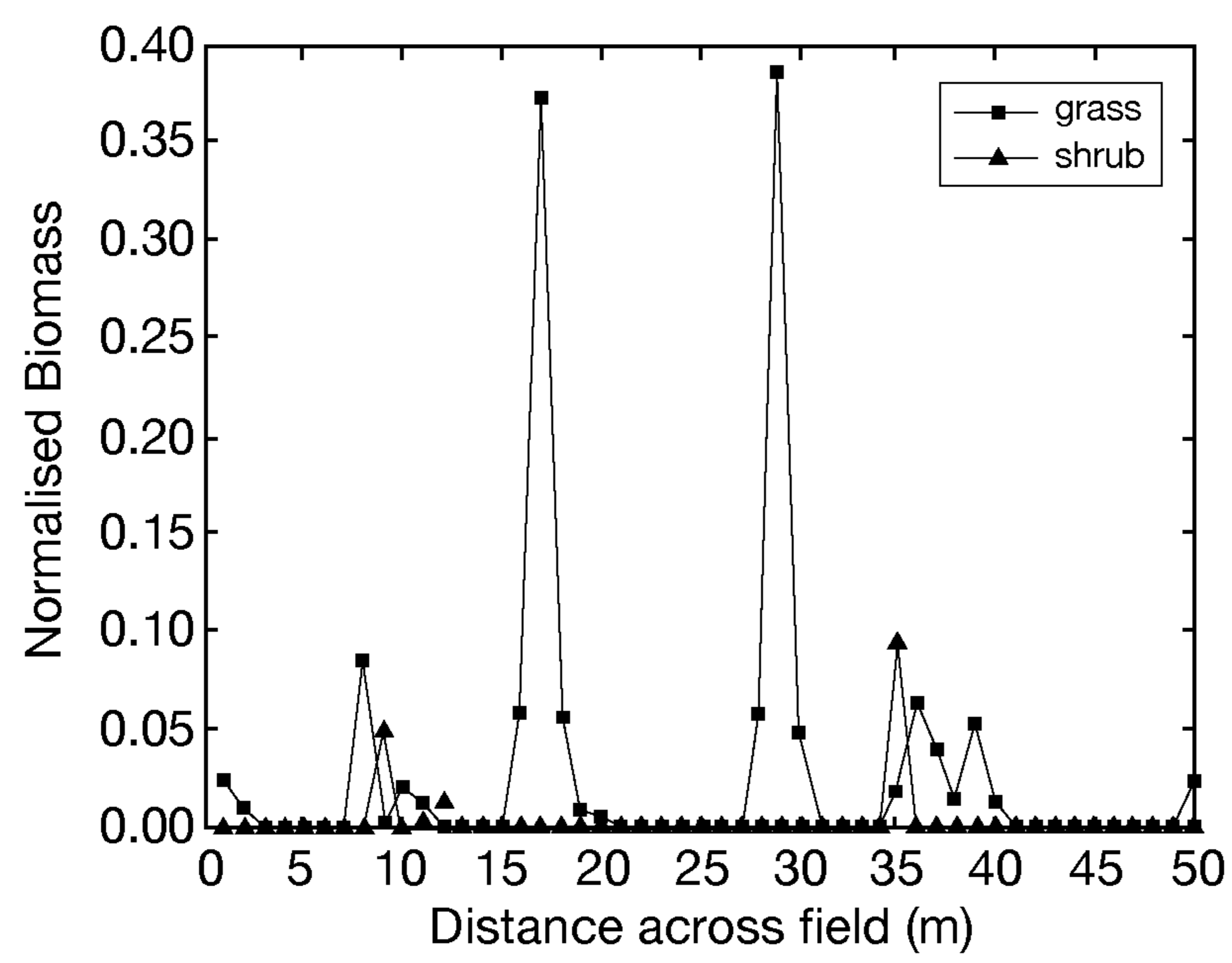
1935



1945



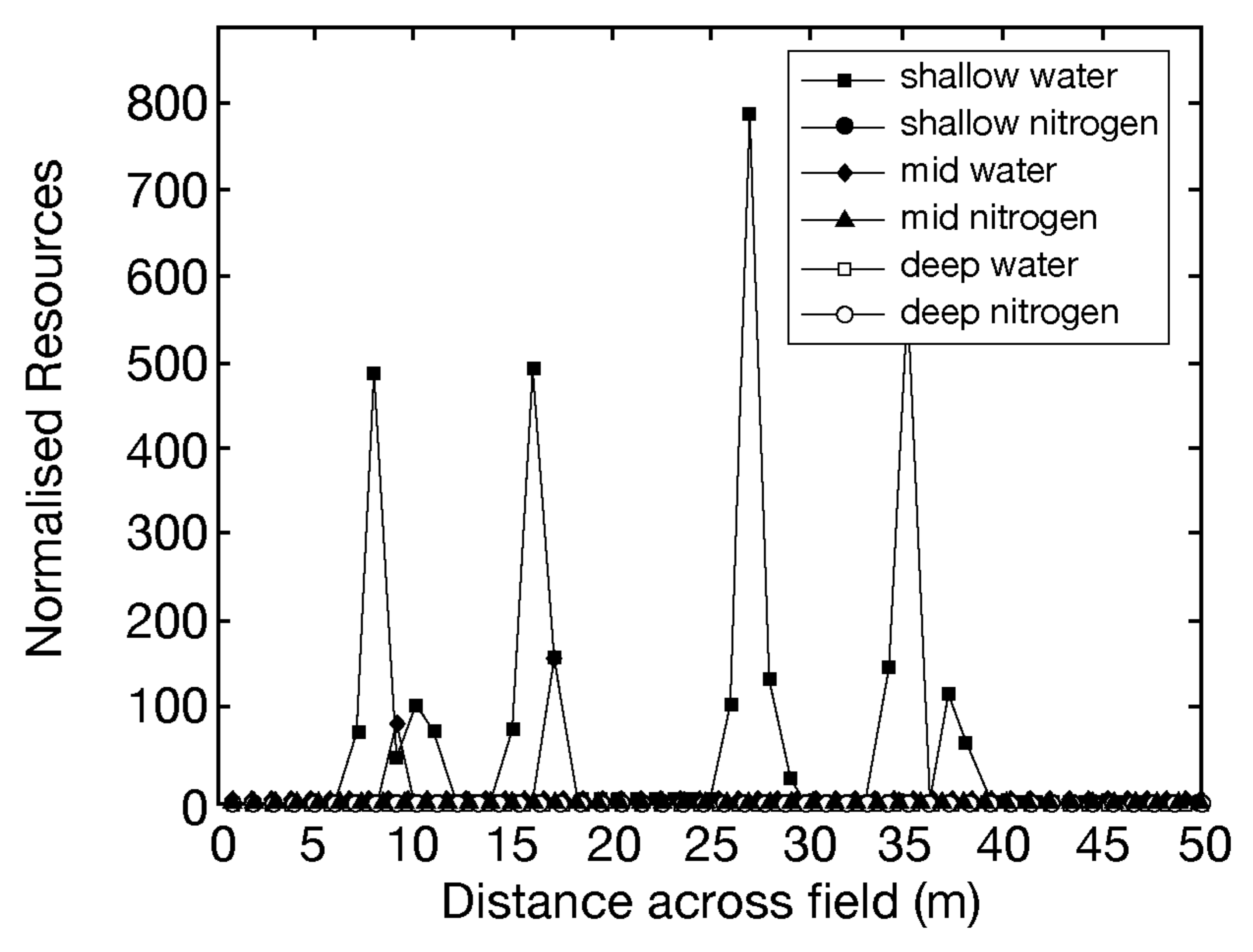
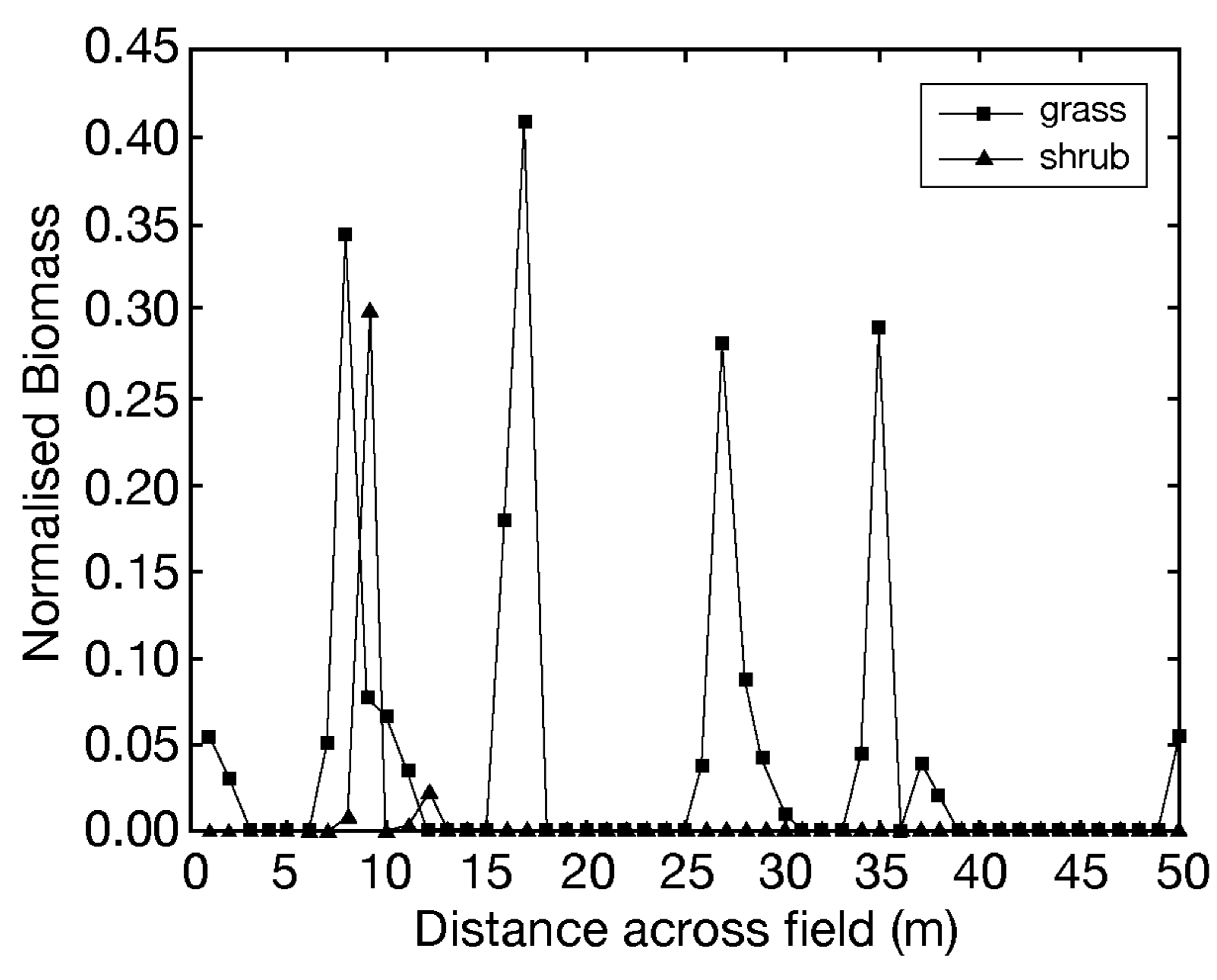
1955



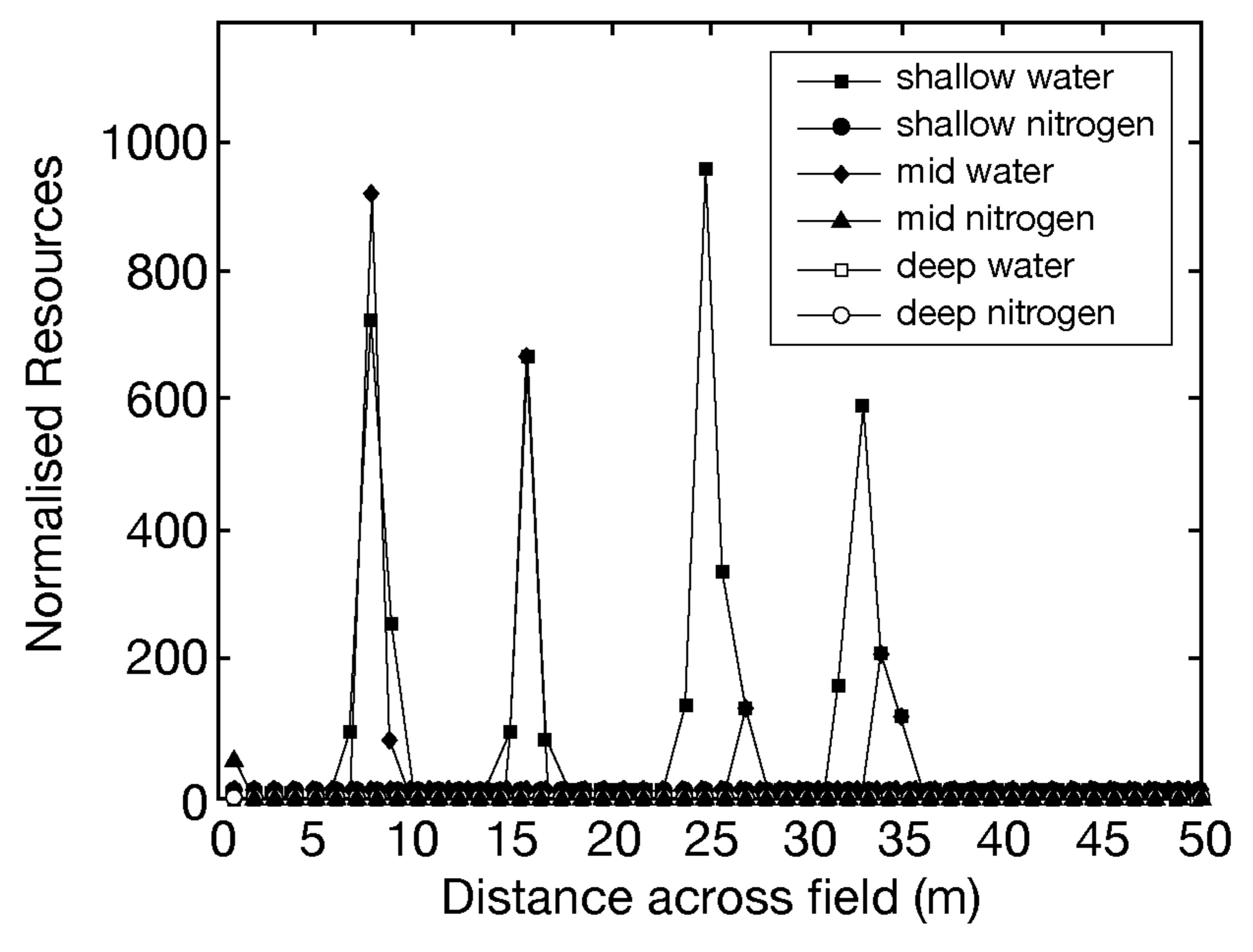
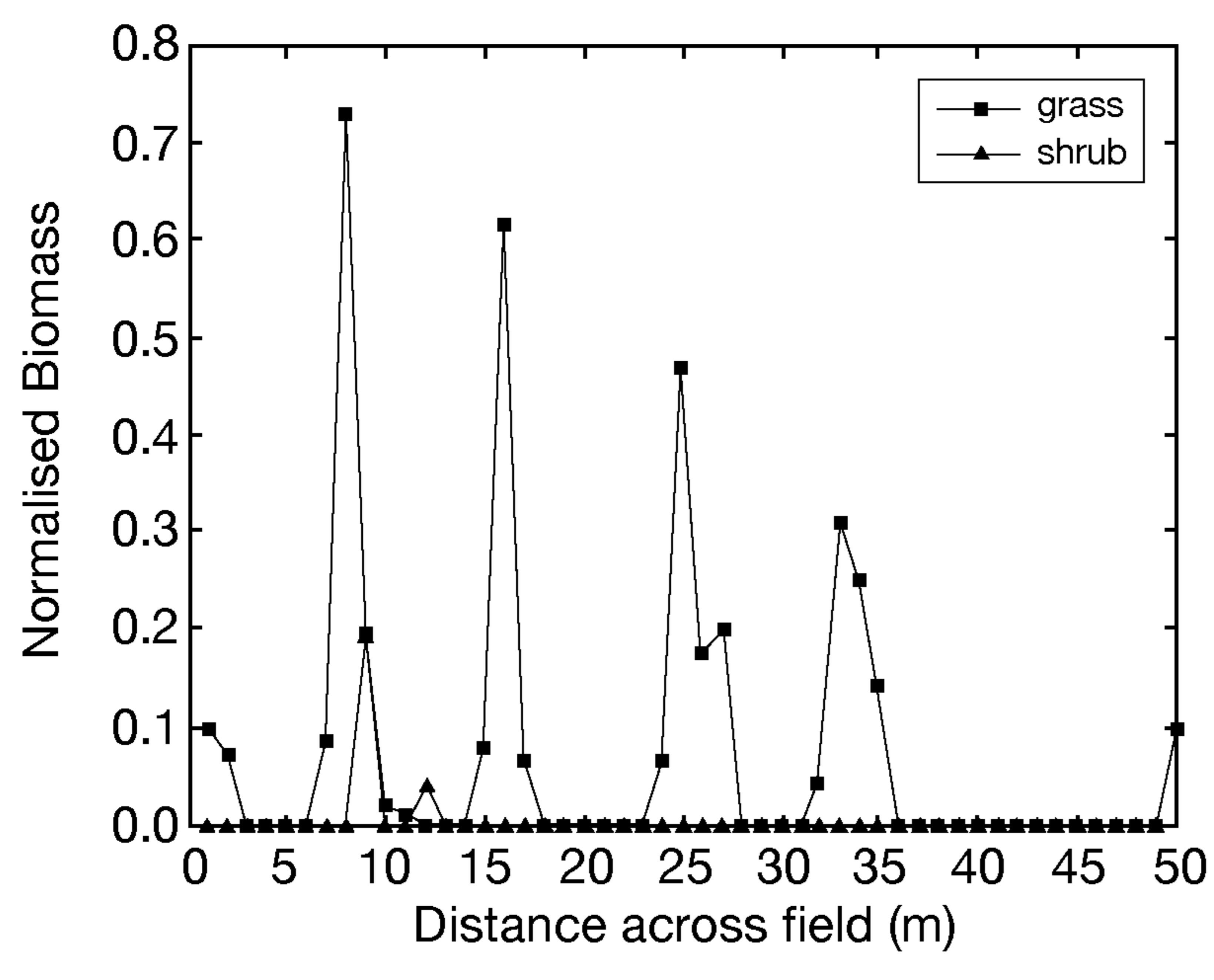
Year **Biomass of grass and shrub**

Resource density in each soil layer

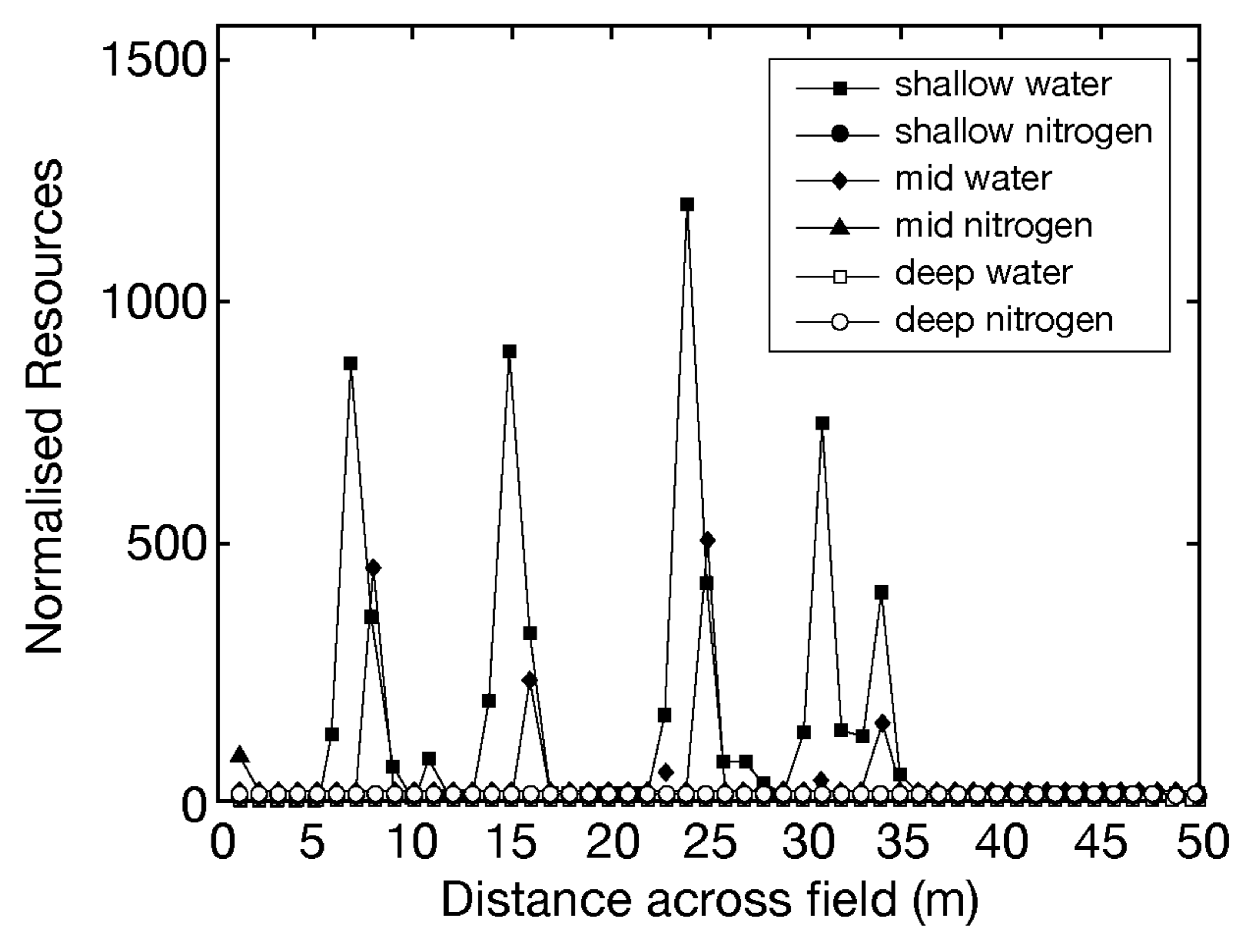
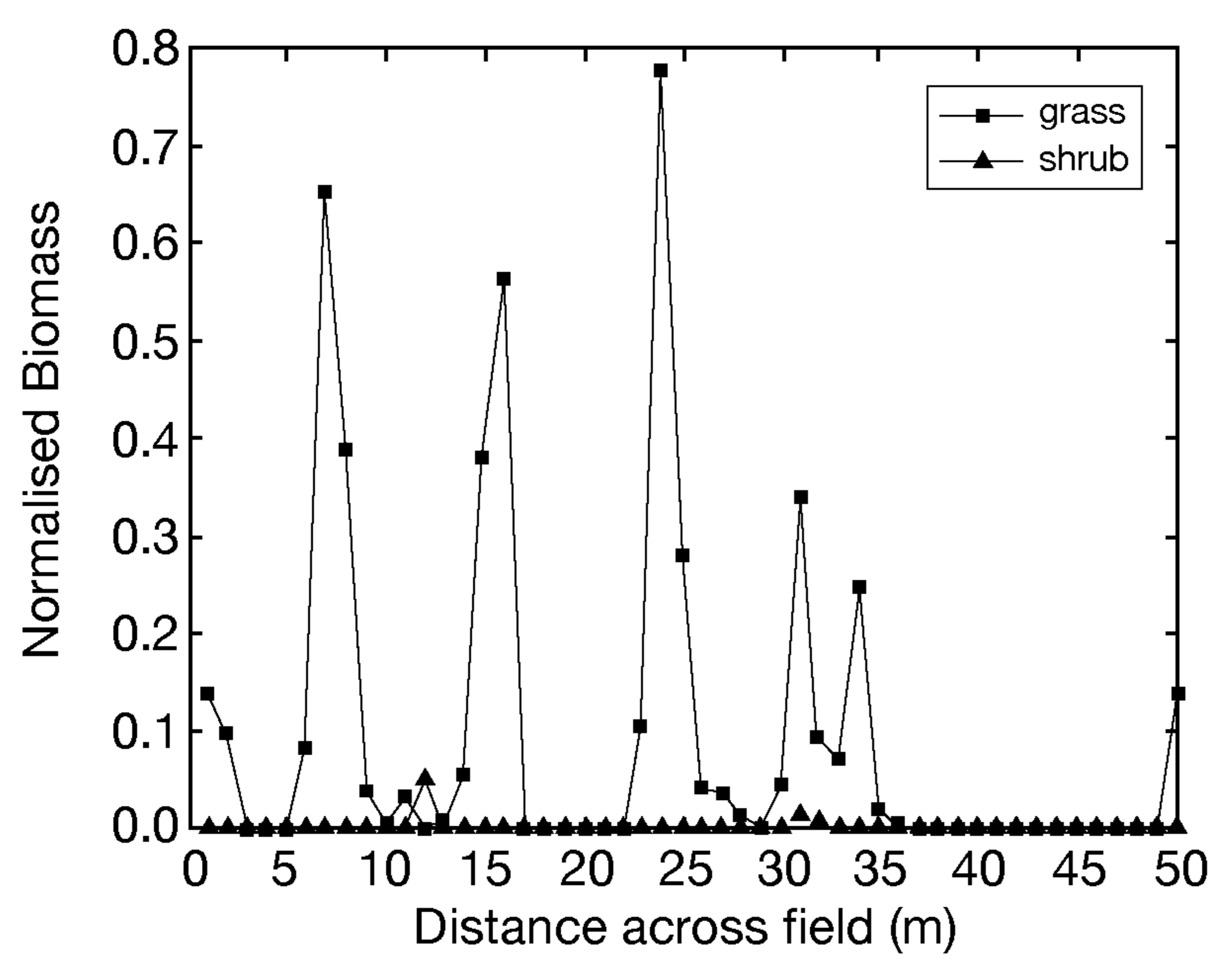
1965



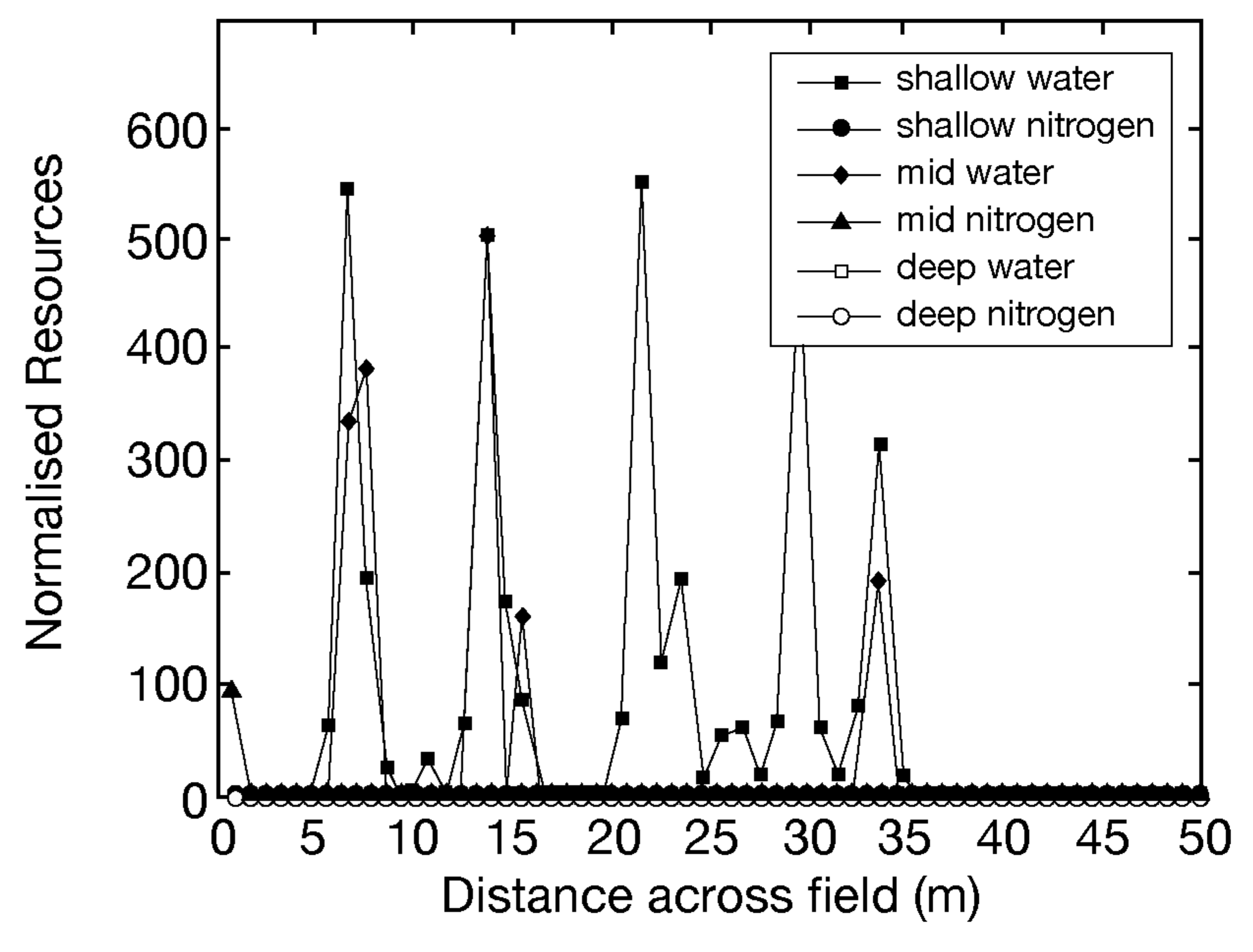
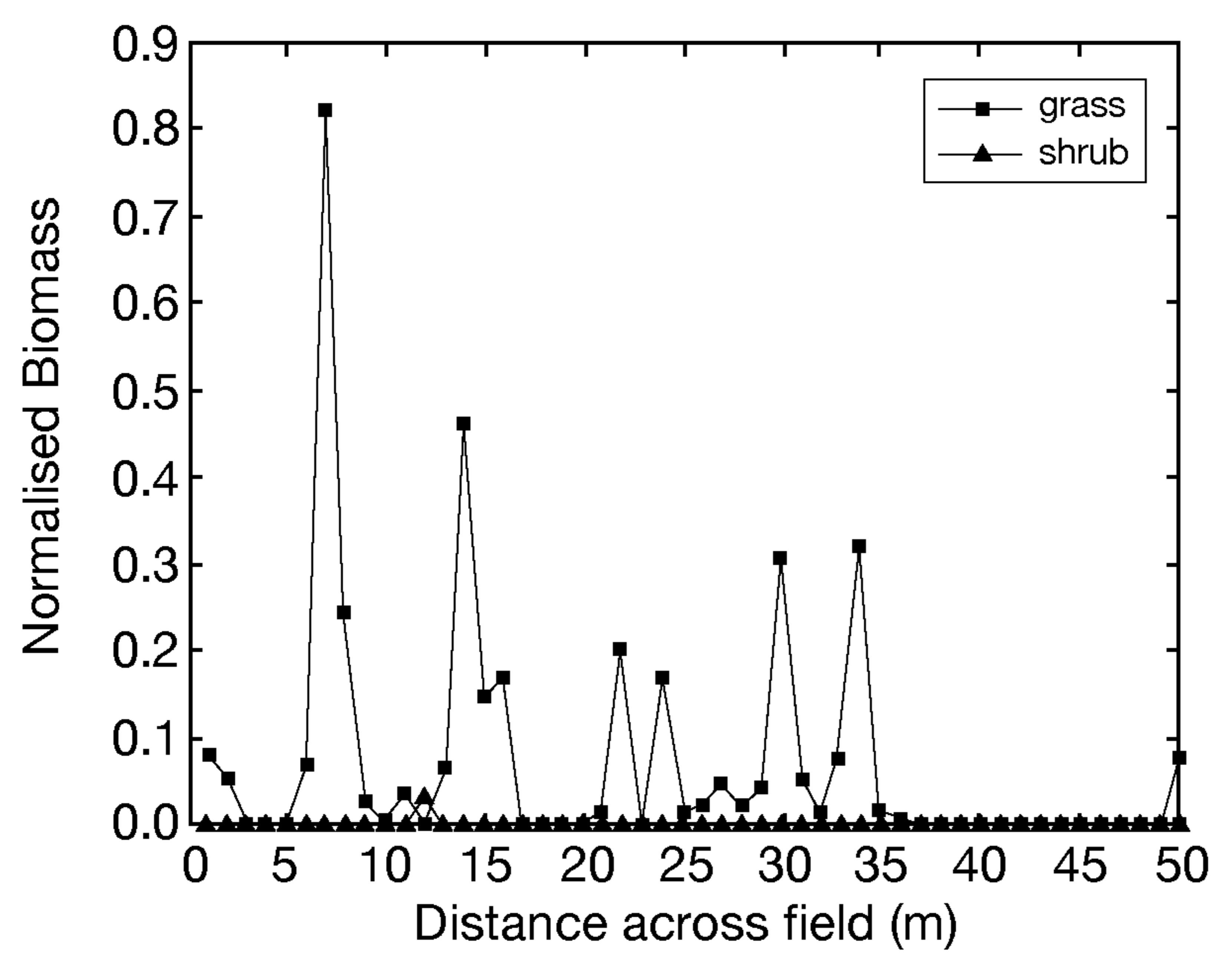
1975



1985

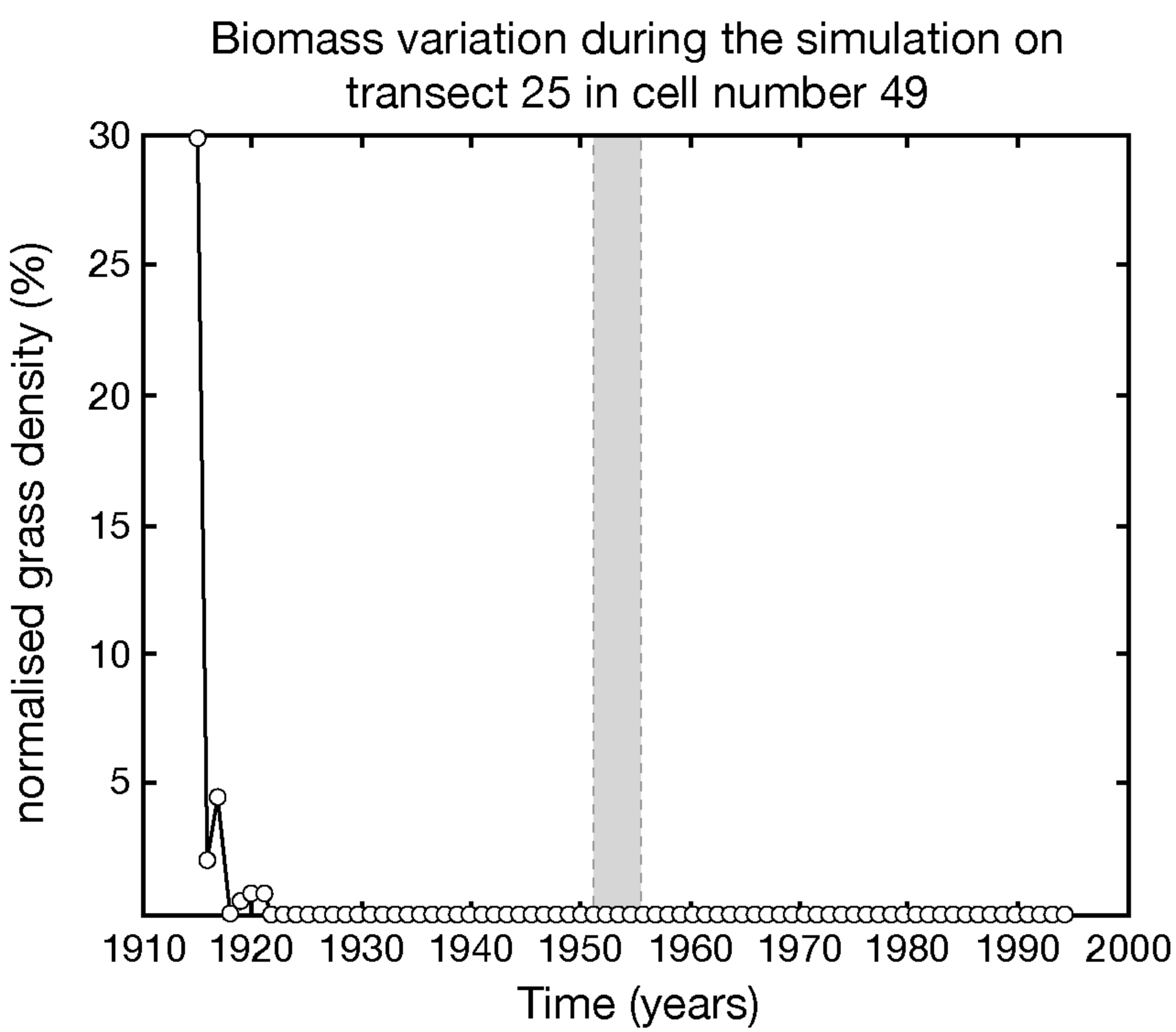


1995

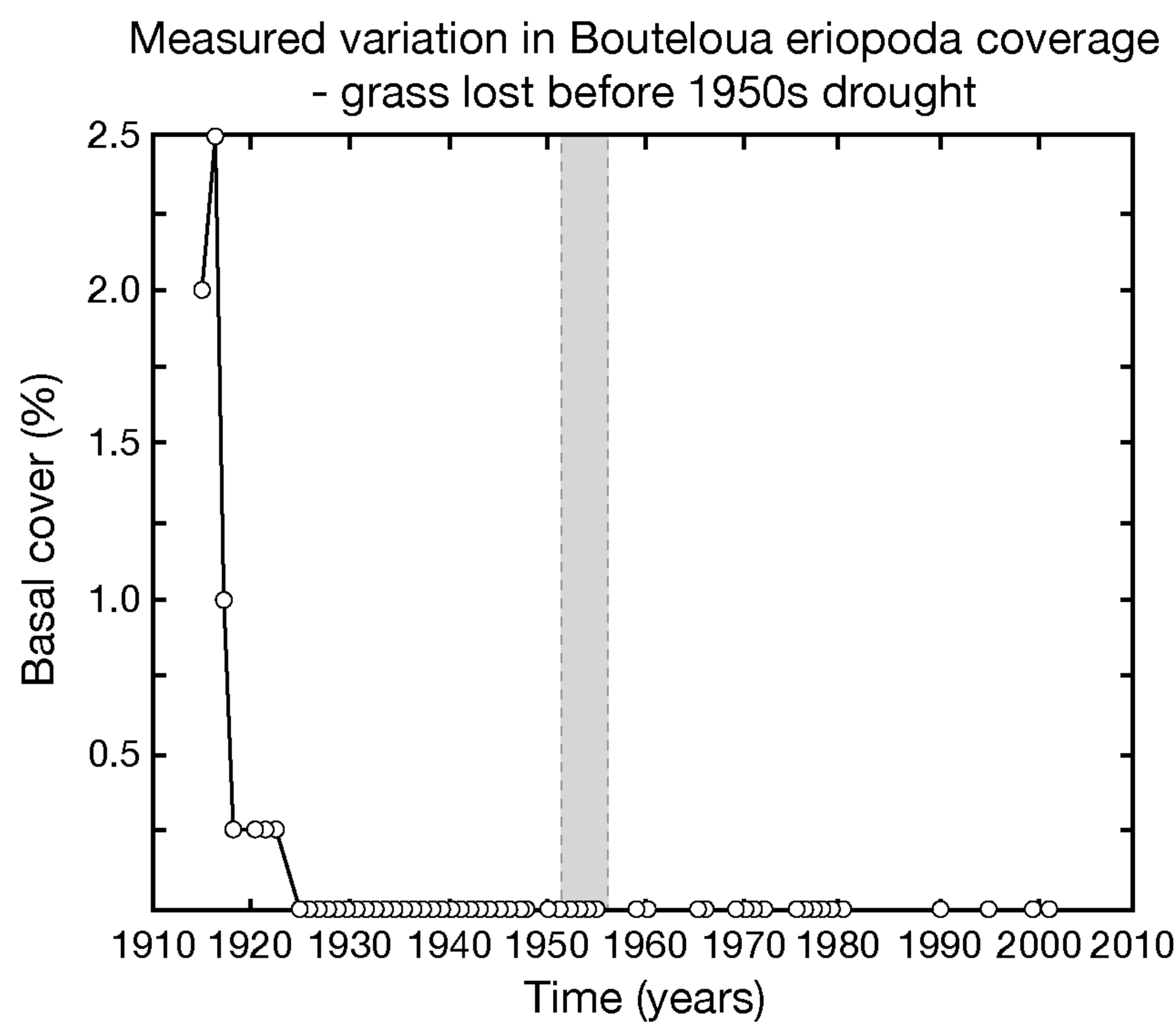


(a) Data illustrating grass biomass lost prior to the 1950s drought

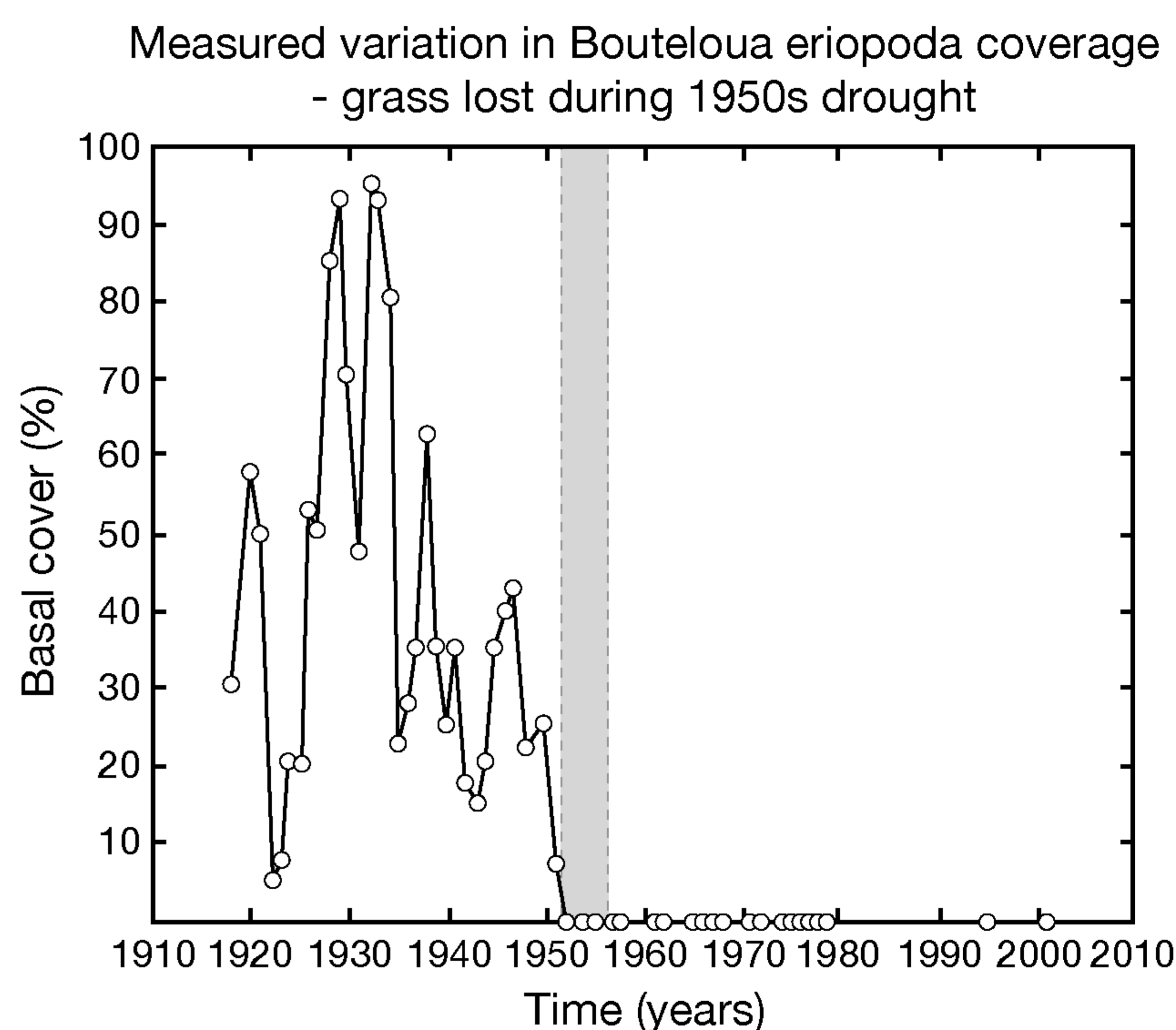
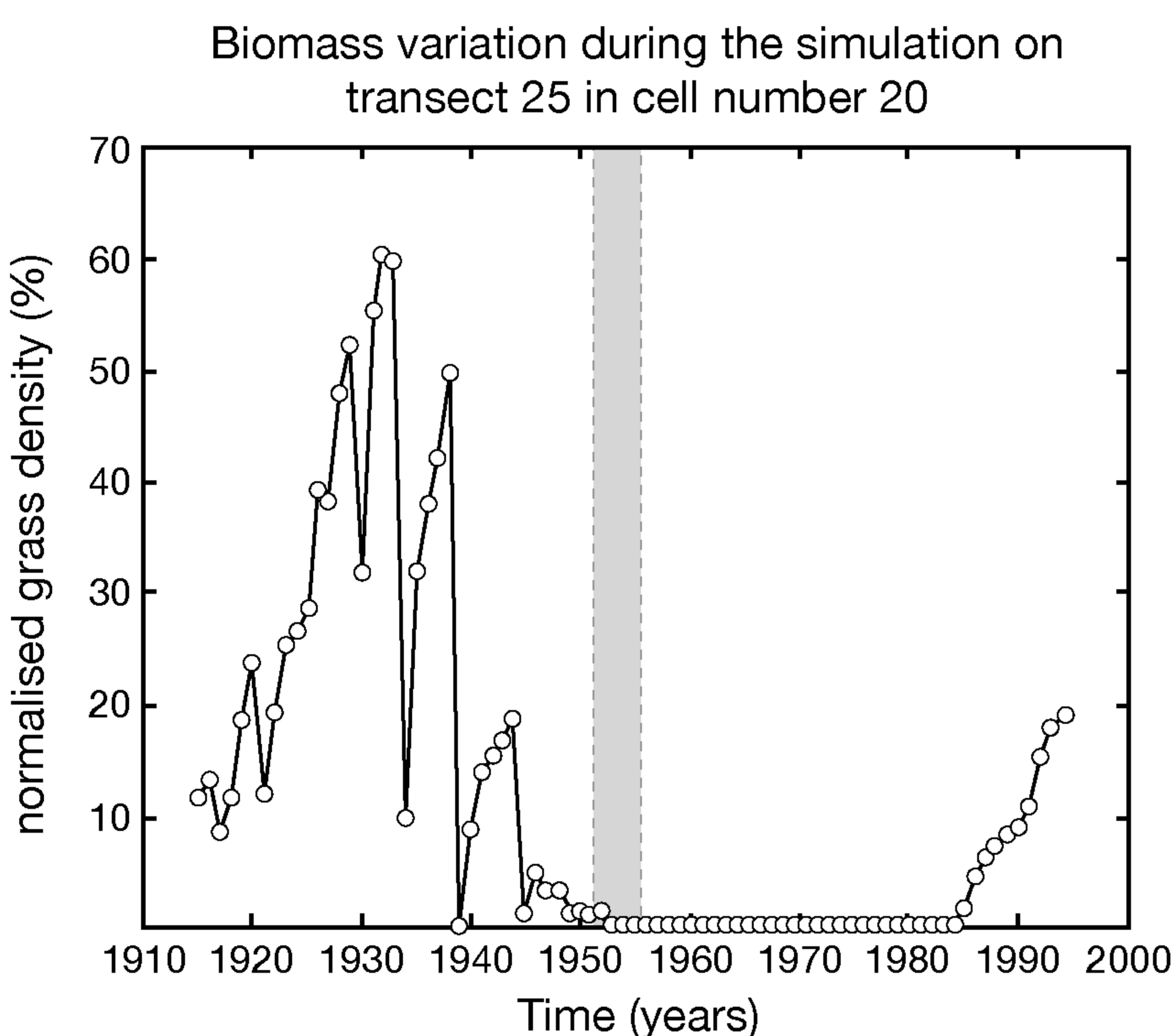
Modelled grass response



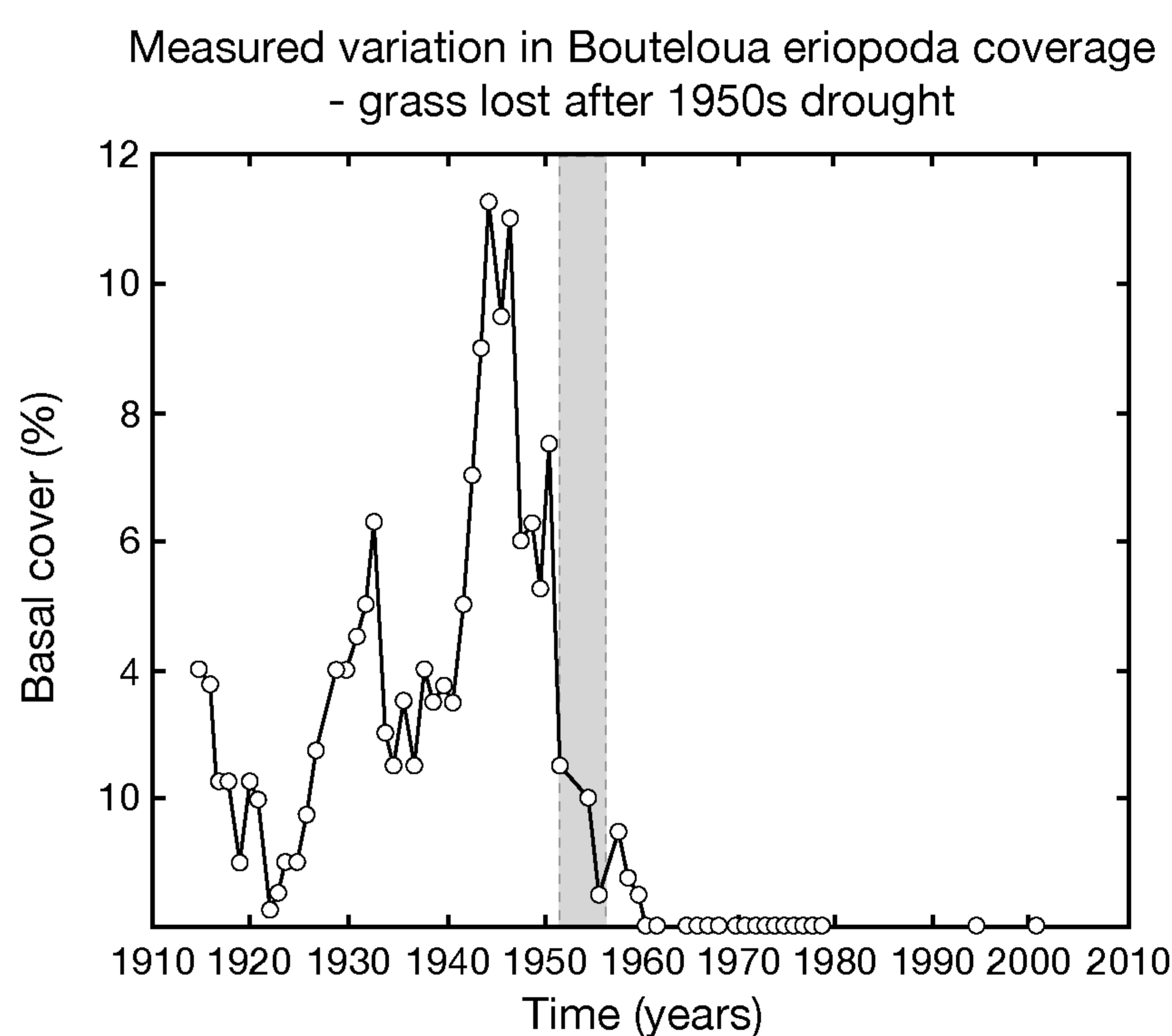
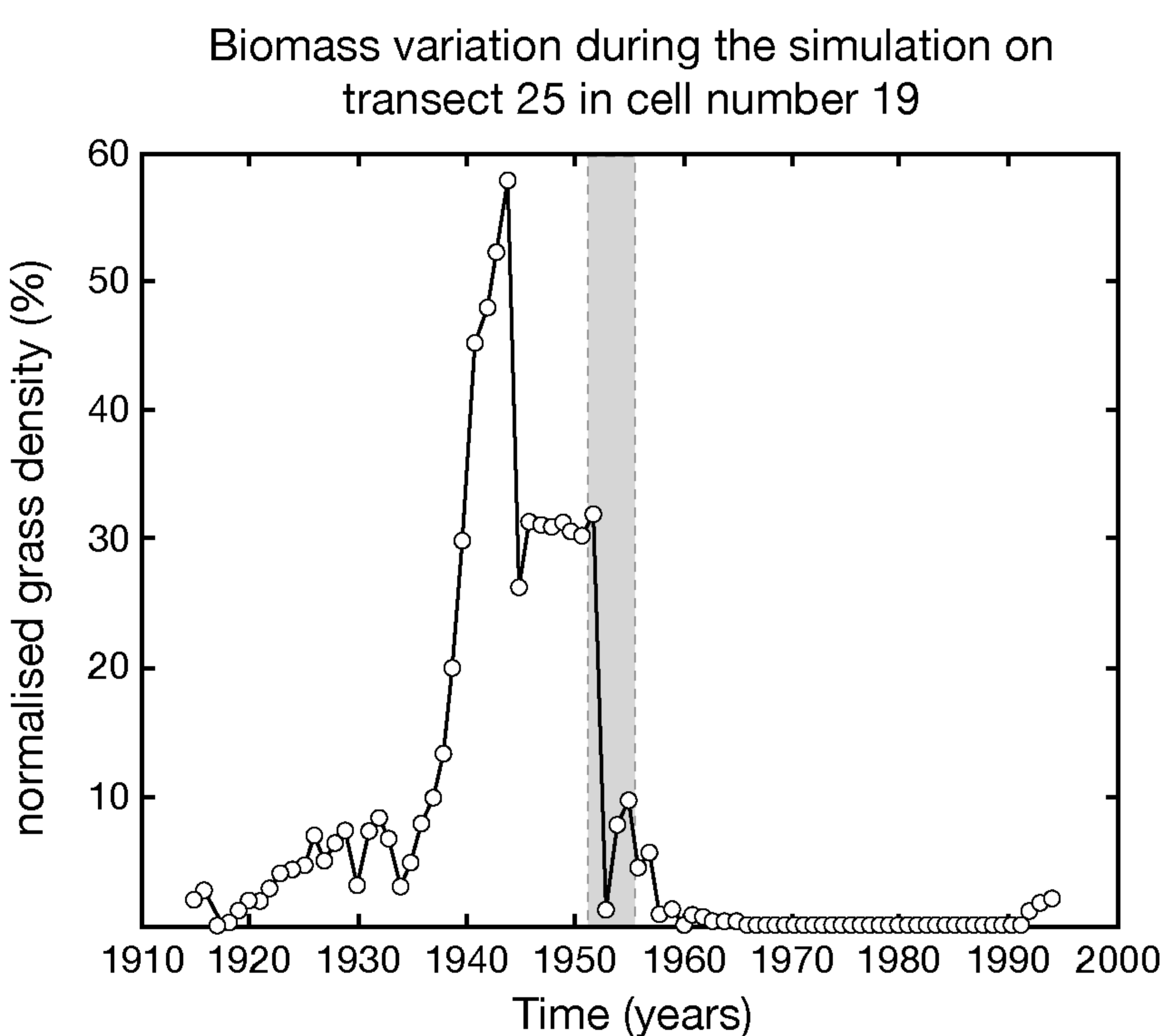
Measured grass response



(b) Data illustrating grass biomass lost during the 1950s drought



(c) Data illustrating grass biomass lost after the 1950s drought



(d) Data illustrating grass biomass that persists after the 1950s drought

