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Regionalization of Flood Data Using Probability Distributions and Their Parameters

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Research Report No. 173

REGIONALIZATION OF FLOOD DATA USING

PROBABILITY DISTRIBUTIONS AND THEIR PARAMETERS

By

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December 1989

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ABSTRACT

The U. S. Geological Survey recently used the method of residuals to delineate seven flood regions for the State of Kentucky. As an alternative approach, the FASTCLUS clustering procedure of the Statistical Analysis System (SAS) is used in this study to delineate five to six cluster regions in conjunction with statistical properties of the AMF series, like the coefficient of variation as estimated using method of L-moments, LCV, the parameters of the EV1 and GEV flood frequency distributions, and the specific mean annual flood, QSP. For both cluster and USGS flood regions, regionalized flood frequency growth curves are developed and their performance evaluated using Monte Carlo simulation techniques. Flood regions are then evaluated and compared using trends in the hydrological characteristics of important variables, performance of the regionalized flood frequency growth curves, discriminant analysis and regression equations relating flood quantiles to watershed physical characteristics. Results show that the cluster regions are more distinguishable in terms of their flood characteristics than the USGS regions. The regionalized flood frequency growth curves of the EV1 and GEV model are more distinct for the cluster regions than the USGS regions, although their performance in terms of bias and RMSE are The standard errors associated with the comparable. regression equations, developed for predicting the EV1 and GEV flood quantiles, are similar for cluster and USGS regions.

Descriptors: Flood * ; flood frequency * ; simulation * ; regionalization * ;

Identifiers:

Cluster Analysis; discriminant analysis; method of residuals; USGS; Kentucky; flood regions.

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CHAPTER 1

INTRODUCTION

Nature, Scope and Objectives

The problem of estimating flood levels for selected frequencies (or return periods) is fundamental to flood control and mitigation studies. This is often accomplished by the use of flood frequency curves developed from systematic flood records at gauged sites in a watershed. However, due to the short or inadequate flow records at these gauges, the predictive ability of such frequency curves is limited. To overcome this problem, regionalized flood frequency curves are developed using pooled data from all gauges located in a hydrologically homogeneous flood The accuracy and reliability of these regionalized region. curves depends to a large extent on the procedures used to delineate flood regions that have similar flood response. Α review of current literature indicates that a limited amount of work has been done in addressing this vital problem of flood regionalization. Furthermore, there exist three distinct methods of regionalization, as described below, that differ fundamentally in the type of approach used.

Method 1: Perform regionalization using specific flood characteristics of original or transformed annual flood data (referred to as response or dependent variables) as recorded at each of the gauged sites. Included in the analysis will be other hydrologic, physical and climatic characteristics of the watershed (referred to as attributes or independent variables affecting flood response) in

which the gauged site is located. After identifying the homogeneous flood regions, a regional flood frequency curve is developed.

- Method 2: An alternative approach to Method 1 above involves the direct use of the underlying probability distribution and its parameters at each of the gauged sites to accomplish flood regionalization. This is done by first performing a flood-frequency analysis using the annual flood series at each of the gauged sites using commonly accepted probability distributions. Select the most suitable distribution and its parameters describing flood response at each gauge. Perform regionalization using the probability distributions and their parameters. In this context, it must be emphasized that gauges within a homogeneous flood region having similar statistical parameters such as the mean, coefficient of variation, and skewness will not necessarily have similar underlying probability law of flood response.
- Method 3: First perform a flood-frequency analysis using annual maximum flood data at each of the gauged sites. A regionalization is then carried out using flood quantiles at selected frequencies (example : the 100-year, 50-year etc. flood levels) as the response variables and other hydrologic, physical and climatic characteristics as the independent variables or attributes.

The three methods of regionalization described above differ in the manner in which they utilize flood data at a gauge. Furthermore, the problem remains as to how different the above methods are in defining homogeneous flood regions.

A secondary objective of the proposed study, therefore, will be to examine this issue in detail.

This study will utilize the systematic flood records available at all the gauges employed by the USGS in deriving the flood frequency curve and its parameters that is most appropriate for a particular gauged site. This information will then be used to classify these sites into homogeneous flood regions. The results from this study will provide a valuable comparison by bringing out the inherent differences in the three methods of flood regionalization described above. In addition, it will examine the most suitable flood probability distribution that can be adopted for the State of Kentucky to accurately describe the flood response of each watershed.

The United States Geological Survey in Louisville is the federal agency primarily responsible for developing regionalized flood information for the State of Kentucky. They have, recently, completed the process of flood regionalization using Method 3 described above. The proposed study will, therefore, examine the problem of regionalization of flood data in the State of Kentucky using Methods 1 and 2. Results from the study should provide a means for comparing these methods of regionalization with an ultimate goal of developing the most accurate and reliable procedure for regionalizing flood data.

With the above discussion in mind, the specific objectives of the proposed study are:

- Perform flood regionalization using Methods 1 and
 2 described above using flood data for the gauges in the State of Kentucky.
- 2) Identify the probability distribution and its parameters that best fits the annual flood series at each of the gauged sites.
- 3) Define homogeneous flood regions based upon the

statistical characteristics of the maximum annual flood data and the probability distributions and their parameters.

4) Compare the homogeneous flood regions as obtained using Methods 1, 2 and 3. Results from the USGS study will be used for Method 3.

Related Research

The index-flood method proposed by the U.S. Geological Survey (Darymple, 1960) is a classic example of early attempts to regionalize flood data. The technique involves the derivation of a regionalized frequency curve using the median values of the ratios of flood discharges at various frequencies to the mean annual flood as defined at each gauge. A major requirement for the application of this method is that the gauges used in the analysis must lie in a hydrologically homogeneous region. Thus, the index-flood method is a convenient way to regionalize flood frequency data provided the hydrologically homogeneous regions are defined a priori.

The use of multiple regression analysis is now a widely accepted method adopted by the U.S. Geological Survey for developing regionalized flood frequency prediction equations (McCabe, 1962; Sauer, 1964; Thomas and Benson, 1970; McCain and Jarrett, 1976; and Richter et al, 1984). The technique involves relating flood characteristics (as reflected by flood magnitudes at selected frequency levels) at a particular gauge to the physiographic, climatic and other variables that affect or control flood response of a watershed. Since this relationship is non-linear, a log transformation is utilized to linearize it. Flood characteristics are obtained at each gauge from log-Pearson Type-III flood frequency distribution (in conjunction with a regionalized coefficient of skewness), the latter derived

using the procedures recommended by the Water Resources Council (U.S. Department of the Interior, 1982). In order to improve the accuracy of these equations, homogeneous regions are defined using the method of residuals. Α residual is the difference between the observed and predicted flood value at a gauged site. This is estimated from the overall regression equation developed for the entire region under investigation. It is assumed that the general trends in these residuals reflect inherent variations in the flood response of various sub-regions. This is the primary basis on which the regionalization of flood data is accomplished. After a detailed analysis of the residuals, a regionalized flood prediction equation is redeveloped using data from all gauges within a homogeneous This method of defining homogeneous sub-regions sub-region. using residuals from an overall regression equation is This is obvious since the causative factors subjective. controlling flood response are not considered explicitly in the process of defining the homogeneous regions. Residuals often reflect statistical variations in the data sample and any trends may be purely incidental. In recognition of this, recent efforts of regionalizing flood data have focussed on the use of more sophisticated statistical For example, DeCoursey and Deal (1974) used methods. discriminant analysis to define homogeneous flood regions using flood and basin characteristics of the watersheds as defined at each streamflow gauge. The basic approach is to classify homogeneous regions using the concepts of cluster analysis. Clusters or groups are formed using flood and basin characteristics at each gauge with the basic premise of maximizing within group similarity while at the same time minimizing between group similarity. A complete linkage algorithm of forming clusters, as proposed by Sokal and Sneath (1963), is used. Discriminant analysis is used to determine any misclassification of points or stations into a cluster or group. Tasker (1982) extended this method of

regionalizing flood data to gaging stations in Arizona. It must be pointed out that, in both cases, flood characteristics were obtained from flood frequency curves as defined at each gauge and hence the general approach is similar to Method 3 described in the previous section.

A completely different approach of regionalization of flood data than the one described above was adopted by Wiltshire (1986) in his efforts to define homogeneous flood regions in England (this is similar to Method 1 described in the previous section). Instead of using flood estimates obtained from a flood frequency curve, his approach incorporates specific properties (statistical) of the flood series as the response variable. An iterative search is then employed using the basin characteristics as the independent variables (or attributes) so as to minimize the variance of the response variable within a cluster or group while simultaneously maximizing the variance between groups. The multivariate technique used in the analysis is referred to as Analysis of Variance (ANOVA) for a single response variable and Multivariate Analysis of Variance (MANOVA) for more than one response variable. The main advantage of Wiltshire's approach is that flood data at a gauge are considered explicitly and the use of a fitted flood frequency curve is avoided. However, as pointed out by Wiltshire (1986), there are two weaknesses to his procedure. The first of these is that the annual maximum flood series at each site is characterized by only one response variable, namely, the coefficient of variation. For example, no consideration is given for other flood characteristics like the coefficient of skewness. The second problem is that the resulting solution in terms of basin groupings may not be unique, i.e. different basin characteristics may also produce a statistically significant result. The latter problem could be resolved to a certain extent using physical reasoning and geographic regions.

The use of probability distribution and its parameters for regionalizing flood data has been attempted by several investigators. Such approaches are similar to Method 2 described in the previous section. Houghton (1977) used the Wakeby distribution and its parameters for regionalizing flood experience in the United States and proposed four such distributions for use in flood prediction. Kuczera (1982) examined the relative performance of the Wakeby distribution in estimating extreme flood events in comparison to other more parsimonious probability distributions. The performance was measured using a mean square criterion. In a parallel study Kuczera (1982) shows how empirical Bayes procedures can be used to combine site-specific and regional information to improve upon site-specific estimators. Rossi et al (1984) regionalized annual flood series using the at-site estimates of a two parameter extreme value probability distribution. Synthetic flood data, generated using Monte Carlo techniques, was used to test the relative performance of several regionalization methods by Lettenmaier and Potter (1985). Their results show that for annual flood series having a high coefficient of variation, improvements in regional flood estimation will come from improved estimators of the at-site mean annual flood, rather than the regional (normalized) flood frequency distribution. An overview of recent efforts in flood regionalization is given by Greis (1983).

Although considerable work, as discussed above, has been advanced in developing robust flood frequency probability distributions, little work has been done in addressing the fundamental question of the selection of homogeneous flood regions. This is an extremely vital step in any effort to regionalize flood data based upon such information from specific gauged sites.

CHAPTER 2

RESEARCH PROCEDURES

The accuracy and precision with which flood levels (particularly those associated with large return periods such as the 100-year flood level), can be estimated at gauged and ungauged streamflow sites is primarily influenced by (Cunane, 1987):

- The form of the underlying flood frequency distribution or model that best describes the underlying law of flood response and the method of estimating its parameters.
- Amount and type of data used: a) at-site data; b)
 at-site/regional and c) regional without at site data.
- 3) Type of flood frequency model: a) Annual maximum (AM) flood series and b) Peaks over threshold (POT) flood series (partial flood series).

The above factors are incorporated into the study procedure as discussed below.

CHOICE OF A FLOOD FREQUENCY MODEL

The choice of a suitable parent probability distribution and the method to estimate its parameters constitutes, by far, the most difficult step in the development of a flood frequency model to best describe the flood response of a watershed. The success of any flood regionalization to estimate flood quantiles accurately is heavily dependent on this choice. The major problem arises from the fact that the true population flood frequency

distribution that best fits the AMF data at a site is and will, at least in the near future, be never known. However. numerous efforts by researchers over the past few decades has led to a general consensus that the annual maximum floods come from populations with positively skewed distributions and that these distributions are relatively thick-tailed. Hence, the focus on contending probability distributions has been primarily on a family of skewed distributions. Furthermore, as suggested by Kuczera (1982), a good flood frequency model must possess the following properties: a) it must have the ability to estimate flood quantiles with least bias and, hence, is efficient (measure of accuracy); b) The model must also be resistant by having the capacity to estimate extreme events, irrespective of which contending distribution best represents the real world, without a disastrous loss of performance as indicated by a suitable measure such as low root mean square error (measure of precision); and c) the flood frequency model must perform well even if a misspecification of the underlying parent probability distribution occurs (a property known as robustness). These are the primary criteria that are given due consideration in the present study for testing the performance and suitability of flood frequency models selected for describing flood experience in -Kentucky.

a) Flood Frequency or Probability Distributions: Numerous probability distributions have been used to fit AMF data. The following is a list of general forms of probability distributions (refer to Table 2.1) that have been used by various investigators either directly or in a simplified form (example: 2-parameter distributions):

- 1. Generalized Extreme Value (GEV) and its special case Extreme Value Type-I (EV1 or Gumbel)
- 2. Generalized Normal and log-Normal

- 3. Pearson and log-Pearson Type-III
- 5. Wakeby
- 6. Generalized Pareto
- 7. Generalized Lambda
- 8. Generalized Logistic
- 7. Kappa

Each of the above distributions require at least three parameters to be estimated which characterize the location, scale and shape of the underlying probability distribution, respectively. They have all been tested by numerous investigators using various procedures for estimating their parameters. Recent studies (Wallis and Wood, 1985, Kuczera, 1982, Lettenmaier et al, 1987, Landwehr et al, 1980) have favored the Generalized Extreme Value, GEV, together with its special case, namely, the Extreme Value Type-I, EV1, (referred to as Gumbel) and the Wakeby, WAK, distributions for modeling AMF data. Furthermore, a relatively new approach called L-moments has been recommended for estimating the parameters of these distributions (Hosking, 1989) over the conventional methods used in the past such as the method of moments and the maximum likelihood method. The method of L-moments is closely linked to the probability weighted moments (PWM) method of estimating parameters as first introduced by Greenwood et al, 1979 and later used by numerous investigators (Landwehr et al, 1979, Landwehr et al, 1980, Hosking et al, 1985, Hosking and Wallis, 1987, Wallis and Wood, 1985, Kuczera, 1982). A brief discussion of the L-moments method, in conjunction with the theory of probability weighted moments (PWM), is given below. This method is chosen as the preferred method for estimating the parameters of the Gumbel(EV1), Generalized Extreme Value (GEV) and the Wakeby (WAK) flood frequency probability distributions in the present study.

Distribution	Code	Number of parameters	Parameters	F(x), x(F)
Generalized extreme-value	GEV	3	ξαk	$F = \exp \left[\{ 1 - \frac{\xi}{2} - \frac{\xi}{2} \}^{1/k} \right]$ x = \xi + \alpha \{ 1 - \left(- \log F)^k \} \/k
Generalized logistic	GLO	3	ξακ	$F = \frac{1}{[1 + (1 - k(x - \xi)/\alpha)^{1/k}]}$ $x = \xi + \alpha [1 - ((1 - F)/F)^{k}]/k$
Generalized Normal	GNO	3	ξ a k	$F = \Phi[-k^{-1}\log(1 - k(x - \xi)/\alpha)]$ x(F) not explicitly defined
Gencralized Pareto	GPA	3	ξak	$F = 1 - \{1 - k(x - \xi)/\alpha\}^{1/k}$ $x = \xi + \alpha \{1 - (1 - l)^k\}/k$
Gumbel	gum	2	ξα	$F = \exp\left[-\exp\left(-\frac{(x-\xi)}{\alpha}\right)\right]$ $x = \xi - \alpha \log(-\log F)$
Kappa	KAP	4	ξakh	$F = [1 - h(1 - k(x - \xi)/\alpha)^{1/k}]^{1/h}$ $x = \xi + \alpha [1 - \{(1 - F^h)/h\}^k]/k$
Wakeby	WAK	5	ξαβγδ	$F(x) \text{ not explicitly defined}$ $x = \xi + \alpha \{1 - (1 - F)^{\beta}\} / \beta - \gamma \{1 - (1 - F)^{-\delta}\} / \beta$

TABLE 2.1. Common Probability Distribution Used in Flood Frequency Analysis (Hosking, 1988)

:

b) Method of Estimating Parameters: Probability Weighted

Moments and L-Moments: A probability distribution, having a distribution function F = F(x) = P(X < x) of a random variable X, may be characterized by probability weighted moments defined as (Greenwood et al, 1979):

$$M_{p,r,s} = E[X^{p} \{F(X)\}^{r} \{1 - F(X)\}^{s}]$$

$$= \int \mathbf{x}^{\mathbf{p}} \{ \mathbf{F}(\mathbf{x}) \}^{\mathbf{r}} \{ 1 - \mathbf{F}(\mathbf{x}) \}^{\mathbf{s}} d\mathbf{F}(\mathbf{x}),$$

(2.1)

$$\int_{0}^{1} \left[\mathbf{x}(F) \right]^{p} F^{r} (1 - F)^{s} dF$$

where p, r and s are real numbers. If r=s=o and p is a non-negative integer then $M_{p,0,0}$ represents the conventional moment of order p as used in the method of moments. If p, r and s are positive integers then the probability weighted moment, $M_{p,r,s}$ can be related to the expected value of the k-order statistic, $X_{k:n}$, of a random sample of size n drawn from the distribution F by the following relationship:

$$M_{p,r,s} = \frac{r!s!}{(r+s+1)!} EX_{r+1:r+s+1}^{p}$$
(2.2)

In particular, the probability weighted moments, $M_{1,0,r'}$ and $M_{1,r,o'}$, which are linear functions of the expected value of the k-order statistic, $X_{k:n}$, are sufficient to characterize a distribution and can be defined as follows:

$$\alpha_{r} = M_{1,0,r} = E[X[1 - F(X)]^{T}], r = 0,1,...,$$

$$= \int_{0}^{1} x(F) [1 - F(x)]^{K} dF \qquad (2.3)$$

$$= \frac{1}{n} \sum_{i=1}^{\Sigma} x_{i} (1 - F_{i,k})^{K}$$

$$\beta_{r} = M_{1,r,0} = E[X[F(X)]^{T}], r = 0,1,...,$$

$$= \int_{0}^{1} x(F) [F(x)]^{K}$$

$$= \frac{1}{n} \sum_{i=1}^{\Sigma} x_{i} (F_{i,k})^{K}$$
(2.4)

Furthermore, α_r and β_r are related by the following equations:

$$\alpha_{r} = \sum_{k=0}^{r} (-1)^{k} {r \choose k} \beta_{k}, \qquad (2.5)$$

$$\beta_{r} = \sum_{k=0}^{r} (-1)^{k} {r \choose k} \alpha_{k}, \qquad (2.6)$$

As stated by Hosking (1986) although the probability weighted moments (PWM's) (Equations 2.3-2.4) can be used to characterize the underlying probability distribution, they are not useful by themselves in defining specific characteristics of a distribution like the scale and shape. Instead, certain linear functions of the PWM's known as L-moments give a better description of the location, scale and shape of a probability distribution. As shown later PWM's and L-moments are closely related.

Consider a real-valued random variable, X, having a distribution function, F(x) and an inverse function, x(F),

and let $X_{1:n} < X_{2:n} < \dots X_{n:n}$ be the ordered statistic of a random sample drawn from the population distribution of the random variable X. L-moments can then be defined as a linear combination the expected value of the above order statistic as (Hosking, 1986):

where, the expected value of an ordered statistic, $E_{j:r}$, is defined as:

$$EX_{j:r} = \frac{r!}{(j-1)!(r-j)!} \int x \{F(x)\}^{j-1} \{1 - F(x)\}^{r-j} dF(x)$$
(2.8)

Substituting Eq. 2.8 in Eq. 2.7, expanding the binomials in F(x) and summing the coefficients of each power of F(x) gives the following final expression that can be used to calculate the L-moments.

$$\lambda_{r} = \int_{0}^{1} x(F) P_{r-1}^{*}(F) dF, r = 1, 2, ..., \qquad (2.9)$$

where,

$$P_{r}^{*}(F) = \sum_{k=0}^{r} p_{r,k}^{*}F^{k}$$
(2.10)

and

$$p_{r,k}^{*} = (-1)^{r-k} {r \choose k} {r+k \choose k}.$$
 (2.11)

Note the similarity of Eq. 2.9 with the PWM as defined in Eq. 2.1. The L-moments are simply linear combination of the PWM's, $M_{1,0,r}$ (Eq. 2.3 and 2.5) and $M_{1,r,0}$ (Eq. 2.4 and 2.6) and, hence, are closely related by the following relationships (Hosking 1986):

$$\lambda_{r+1} = (-1)^{r} \sum_{k=0}^{r} p_{r,k}^{*} \alpha_{k} = \sum_{k=0}^{r} p_{r,k}^{*} \beta_{k}, r = 0, 1, \dots$$
(2.12)

c) Interpretation and estimation of L-moments: As pointed out by Hosking (1989), the L-moments λ_1 , λ_2 , λ_3 , \dots , λ_r , and L-moment ratios $\tau_3 = \lambda_{3\lambda_2}, \tau_4 = \lambda_{4/\lambda_2}, \dots, \tau_r = \lambda_{r/\lambda_2}$ are useful quantities for summarizing a probability distribution. The L-moments are similar to conventional central moments while the L-moment ratios are similar to the conventional moment ratios. The first L-moment, λ_1 , is equal to the mean and is, therefore, regarded as a measure of the central tendency or location, the second L-moment, λ_{s} , is a measure of scale or dispersion like the variance or standard deviation. The moment ratios, τ_1 and τ_2 , which are dimensionless forms of the third and fourth L-moments (λ_3 and λ_{L}), are measures of skewness and kurtosis, respectively. Thus, these L-moments and ratios together are sufficient to estimate parameters that describe the location, scale, skewness and kurtosis of a flood frequency

distribution . Higher order L-moments and ratios have similar interpretation as conventional method of moments for further describing the character of the underlying probability distribution.

The L-moments described above must be estimated from observed maximum annual flood data at a gauged site prior to any flood regionalization effort. A natural estimator of each L-moment (refer to Eq. 2.7 above) based on an observed sample of data is a linear combination of the ordered data values. Such an estimator is known as an L-statistic. In practice, therefore, the L-moments can be estimated from an ordered (lowest to highest value) random sample drawn from an unknown probability distribution. Hosking (1989) presents two such estimation procedures. The one used in this study is referred to as a plotting position estimator. A plotting position, p_{i:n}, is a distribution-free estimator of the probability of non-exceedance, F(x_{i:n}), of an ordered random variable X_{i:n}. Although this estimator is biased, Hosking (1989) has observed in his study that it gives good estimates of the parameters and quantiles when a distribution is fitted to the data. In particular, Hosking (1989) concludes that the plotting position estimator of the form $p_{i:n} = (i - 0.35)/n$, where i is order number (or rank) of observed data value, x_{i:n}, of random variable X_{i:n}, and n is the sample size, gave good results for generalized extreme value distribution. Thus, the following equations are used in the study to estimate the L-moments and PWM's, α_r (refer to equations 2.9, 2.3 and 2.4, β and respectively).

$$a_{r}[\gamma,\delta] = n^{-1} \sum_{i=1}^{n} (1 - p_{i:n})^{r} x_{i},$$

$$b_{r}[\gamma,\delta] = n^{-1} \sum_{i=1}^{n} p_{i:n}^{r} x_{i},$$

$$l_{r}[\gamma,\delta] = n^{-1} \sum_{i+1}^{n} p_{r-1}^{\star}(p_{i:n}) x_{i}$$
(2.13)

PROCEDURE FOR FLOOD REGIONALIZATION

Statistical estimates of flood quantiles, based on at-site data only, are highly variable due to modeling and sampling error. Consequently, a process of flood regionalization, whereby flood data from several sites within a homogeneous flood region (defined a priori) are pooled together, is usually recommended. In the present study, an at-site/regional flood data (refer to 1(b) above) approach is adopted using the historical AMF series (systematic record) (refer to 2(a) above) from each of the gauged sites in Kentucky. This data are transformed to a dimensionless form by dividing each observed flood value by the mean annual flood at that site. An index-flood approach for flood regionalization similar to the one used by Hosking and Wallis (1987), and as described below, was used to pool flood data from gauged sites within a homogeneous flood region (the procedure used to delineate such regions is discussed later). An IBM supplied computer program (Hosking, 1988) is used, with some modifications, to accomplish the regionalization. This computer program allows the development of a regionalized flood frequency distribution for commonly used probability distributions. The method of L-moments is used to estimate regionalized parameters. A step-by-step procedure of the index-flood method used in this study is as follows:

1. Define flood regions that have similar underlying flood response. These regions can be delineated either using the statistical moments required to characterize the underlying parent probability distribution or the parameters of this distribution as estimated using the statistical moments. In either case, the basic premise is that regions having similar statistical moments or parameters of the probability distribution must be homogeneous with respect to their flood response.

Alternatively, such regions may be delineated on the basis of the physical characteristics of the watersheds that control flood response. In this study, both approaches are used to identify flood regions.

- 2. Within each region assume that the regional flood quantile estimate for a given return period, T, is given by q_T . This estimate is derived from the probability distribution of normalized flood data (dimensionless flood variate) and hence is scale independent. The normalized flood variate, X is obtained by dividing each flood observation, Q_i , at a site by an index flood, Q_I . The latter is usually taken as the mean annual flood at the site as is done in this study.
- 3. Estimate the at-site mean annual flood, Q, at each site, i, within a region using the average of observed raw flood data as required in step 2 above and the following step.
- 4. Combine estimates, q_{T} , and \overline{Q} , to obtain the flood quantile estimate, $Q_{T_{i}}$ at site i within the region.
- 5. The accuracy and precision of the flood quantile estimate, Q_{Ti}, at site i is then evaluated using Monte Carlo simulation techniques.

DEVELOPMENT OF REGIONALIZED FLOOD FREQUENCY GROWTH CURVES

A frequency growth curve is simply a plot of a cumulative probability density function and can, therefore, be used to compute flood levels at various probability levels of non-exceedance (flood quantiles). In this study these curves are plotted with the normalized flood levels (random variate) on the vertical axis and the probability of non-exceedance on the horizontal axis of a Generalized Extreme Value probability paper. A high value of coefficient of variation and skew prevalent in the AMF data

would cause this growth curve to be steeper reflecting more variability in the data. Thus, a given normalized discharge level will be associated with a smaller return period (or probability of exceedance) than a flatter curve. Furthermore, these growth curves can be directly related to the flood response of the watershed (Acreman and Sinclair, For example, larger watersheds, responding to floods 1986). generated from various sub-watershed contributions, may exhibit greater variability in their flood response than smaller watersheds. Hence, flood data from larger watersheds would have a larger coefficient of variation resulting in a steep growth curve. The shape of the growth is, also, influenced by other watershed physical and climatic characteristics like watershed size, slope, landuse, soil and spatial and temporal effects of rainfall inputs. In any event, differences in the shape of these growth curves (regionalized) do reflect variations in flood response, and can, therefore, be used to assess the degree of heterogeneity of flood response between flood regions.

DELINEATION OF FLOOD REGIONS: CLUSTER ANALYSIS

The purpose of cluster analysis in the context of flood regionalization, is to place gauged sites into clusters or groups such that all the gauges within a cluster have similar flood response and those in different groups have dissimilar flood response. Therefore, the success of any clustering technique would greatly depend on the variables used to define similarity of flood response and some sort of measure to cluster gauged sites that are closer than others with respect to these variables. Since the flood response of any watershed is dependent on the underlying probability law of flood response, it is appropriate to use the statistical moments that characterize this distribution and/or the the parameters of the probability distribution

(as estimated from the moments) as the variables to measure similarity (referred to as response variables in this study) of gauged sites within a cluster. In order to accomplish this, a criterion to group gauged sites having similar statistical moments or parameters (or response variables) is required. A commonly used method is based on the concept of Euclidean distance. In particular the Mahalanhois distance, as defined in the following equation, has the added advantage when compared to an ordinary Euclidean measure since it explicitly accounts for any correlations that might exist between the variables used in clustering.

$$D^{2} = (X_{i} - X_{j})' S^{-1} (X_{i} - X_{j})$$
 (2.14)

where,

= Euclidean distance,

S = pooled within-group covariance matrix.

In this study a clustering technique based on the Euclidean distance measure described above is used to group (or to bring together) gauged sites into homogeneous flood regions or clusters. Several clustering algorithms, such as the average linkage, nearest centroid sorting (referred to as FASTCLUS), complete linkage or Ward's minimum variance can be used to perform the clustering based upon the Eulidean distance given by Equation 2.14 (SAS, 1985). The choice amongst these will depend on the data being analyzed although the FASTCLUS disjoint clustering algorithm has an intuitive appeal over the other methods since its procedure allows for the movement of observations at every step of the clustering process.

a) Choice of the Clustering Algorithm: As mentioned above there are several algorithms that are commonly used in performing cluster analysis. The principal difference between each of these algorithms stems from the manner in which they compute the Euclidean distance measure and the manner in which the clustering is performed. Consequently, the nature of the clusters formed will depend heavily on the variables and their corresponding values. A brief discription of characteristics and biases of the more frequently used clustering algorithms that makes each different or distinct from others is presented in SAS, 1985. These inherent differences are used in this study to make the final choice of the algorithm.

The FASTCLUS clustering technique, as available in the Statistical Analysis System computer software, SAS (SAS Institute 1985), is used to group (or to bring together) gauged sites into distinct flood regions or clusters. This procedure performs disjoint clustering on the basis of Euclidean distances computed from the clustering variables The FASTCLUS procedure differs from hierarchical used. clustering procedures, such as Ward's, by using cluster Initial cluster seeds are observations which are seeds. separated by at least a specified minimum distance. FASTCLUS is an iterative procedure in which cluster seeds are recomputed for each iteration. In each iteration, all observations are assigned to the nearest seed, forming the specified number of clusters, and the seeds are recomputed as the means of the clusters. Observations are then considered as seed replacements using two tests based upon maximizing the distance between seeds. This iteration process continues until a convergence criterion, based upon the maximum distance any seed is changed, is met. Then the final clusters are formed by assigning each observation to the nearest seed. The FASTCLUS procedure is sensitive to outliers.

The FASTCLUS procedure described above is similar to the procedure used by Wiltshire (1986) in his efforts to regionalize flood data in England. In favor of this form of clustering, Wiltshire (1986) points out that "partitioning imposes a certain degree of structure on the data and avoids the undesirable tendency of hierarchical schemes to produce one large dominant cluster located at the centroid of the data with small satellite clusters toward the margins of the data space". A similar situation was observed by the authors when using hierarchical clustering algorithms such as Ward's. This was the primary reason why the FASTCLUS procedure is selected over the other methods in this study. Based on the flood response variables, namely statistical moments required to characterize the underlying probability distribution, the parameters of the probability distribution and the specific mean annual flood, QSP (clustering variables), disjoint clusters or flood regions are successfully delineated.

One of the most difficult problems in cluster analysis is the identification of the optimal number of clusters in a data set that can be clearly distinguished from each other. A review of current literature suggests that several procedures, referred to as stopping rules, available for addressing this vital issue. Such rules are often applied in a subjective manner. To use these rules in the classical "test of hypothesis" setting requires the specification of a null and alternate hypothesis, such as that the data are a random sample from a multivariate normal population. However, it has been shown that there can be large errors associated with these tests if the hypotheses are not stated correctly. Futhermore, there is the additional problem of determining the sampling distribution of the criterion used in the hypothesis testing. Ordinary tests like ANOVA F and t-test are not valid for testing difference between clusters, since clustering methods tend to maximize the separation between clusters and hence violate the basic

assumptions of such tests. In view of this, formal tests of hypothesis are not used in this study. Instead, a number of stopping rules are incorporated in a subjective manner while selecting the optimum number of clusters. In doing so, the principal objective of identifying homogeneous cluster or flood regions that can be discriminated easily based upon the attribute variables is given primary emphasis. Milligan and Cooper (1983) used Monte Carlo simulations to evaluate the performance of 30 stopping rules commonly used in cluster analysis. Amongst these, several rules which gave good performance are selected and discussed below. Furthermore, since only 253 gauged sites are being used in the flood regionalization study, it seemed impractical and physically unrealistic to examine more than ten clusters. Cosequently, the following stopping rules, as presented by Milligan and Cooper (1983), are applied to 10 or fewer cluster regions.

- 1) The goodness of fit criterion, R^2 , has the usual interpretation of the proportion of variance accounted for by the clusters. Ward's algorithm attempts to maximize this when deciding on the clusters to merge at each stage of clustering. As clusters are merged R^2 will decrease and, hence, a rule of thumb is to stop clustering whenever there is a significant drop in the value of this criterion.
- 2) The ratio criterion is defined as the ratio of within cluster sum of squared errors when the data are split into two clusters to the squared errors when only one cluster is used. In general, small ratio of this criterion leads to the regection of the hypothesis of one cluster. This criterion, as first proposed by Duda and Hart (1973), gave the best performance amongst all the other rules examined by Milligan and Cooper (1984). The ratio criterion can be applied at each stage of the clustering to the subpopulations involved. Thus, at

any stage, if the ratio is small, the two clusters being merged should remain separate. In contrast, a larger value of this ratio would support the collapsing of the two clusters into one. The Duda and Hart ratio criterion can be related to the pseudo- t^2 statistic available in the SAS package by a reciprocal relationship (SAS, 1985). The pseudo- t^2 statistic is a measure of the separation between clusters most recently merged. Thus, a rule of thumb while selecting the optimum number of clusters is to look for small values of this statistic.

- 3) Another stopping rule that performed in the top one-third of the stopping rules studied by Milligan and Cooper (1983), was the pseudo-F statistic. While similar to the F-statistic in ANOVA, the assumptions associated with analysis of variance are not met in the clustering setting and, hence the name "pseudo-F". This statistic provides the measure of separation among all clusters at any step in the clustering process. Ideally, as the number of clusters decreases, the pseudo-F statistic will decrease, then rise at the point where the optimum number of clusters occur, and then fall again (this is referred to as an "elbow" effect). If such is not the case, the pseudo-F will continue to decrease as the clusters are collapsed. In this case, one could look for the largest gap of this statistic in selecting the optimum number of clusters.
- 4) The cubic clustering criterion (CCC criterion) developed by Searle (SAS, 1985) performed as well as the pseudo-F statistic in the simulation runs by Milligan and Cooper(1983). This criterion is a function of the observed R^2 (refer to stopping rule 1) and the expected R^2 assuming that the clusters, as obtained from a uniform distribution on a hyperbox, are hypercubes of the same size. Guidelines for using CCC criterion include the plotting of CCC statistic versus

the number of clusters with the peaks indicating the possible cutoff point for extracting the optimum number of clusters. Peaks associated with a CCC value greater than or equal to 2 indicate a good number of clusters.

b) Selection of Suitable Flood Characteristics for

As stated in the previous section, the success Clustering: of using cluster analysis to delineate homogeneous flood regions depends to a large extent on the variables used to define the flood characteristics at each of the gauged sites (response variables). Since any data set of observations can form clusters, it is imperative to choose variables that reflect the flood experience as accurately as possible in order to ensure flood homogeneity within a cluster. The flood response of a watershed, as measured using the AMF series at a gauge, is stochastic and is, therefore, governed by an underlying probability law (distribution unknown a The latter can be evaluated by fitting the AMF priori). series to an assumed probability distribution using statistical moments of various orders such as the fist order moment. Consequently, it can be postulated that any two gauged sites will have similar flood response if their underlying probability distribution is the same. This would also imply that the statistical moments used to fit the probability distribution (involving the evaluation of its parameters) and/or its parameters must be identical except for the effects of scale. Based upon this premise, the following clustering variables (flood response variables) are initially used to perform cluster analysis using FASTCLUS clustering algorithm. All the clustering variables are standardized prior to clustering in order to suppress any disproportionate effects during clustering.

 L-moment ratios (dimensionless ratios of L-moments) of normalized maximum annual peak flow data from each gauged sites, namely, coefficient of variation, LCV,

coefficient of skewness, LSK and coefficient of kurtosis, LKUR. All these variables characterize the form of the underlying probability distribution. For two-parameter distributions the first L-moment, LCV is adequate while for probability distributions with more than two parameters higher order L-moment ratios will be required.

- 2) The specific mean annual flood, QSP, defined as the ratio of the mean annual flood at each site (as estimated using raw flood data) to the watershed size in square miles.
- 3) The parameters (as estimated using L-moments) of the selected flood probability distributions. The number of parameters used will depend on the distribution selected. Generally, two to three parameters reflecting the location, scale and shape are required for most probability distributions.

The final choice of suitable response variables for obtaining the clusters is based upon the results of cluster analysis, specifically, the ability to extract optimum number of clusters using a cutoff criterion, detailed examination of trends in important hydrological characteristics within and between regions, flood frequency growth curves, discriminant analysis using the attribute variables at each gauged site, and regression analysis relating selected flood levels to watershed physical characteristics. These results are presented in the next chapter.

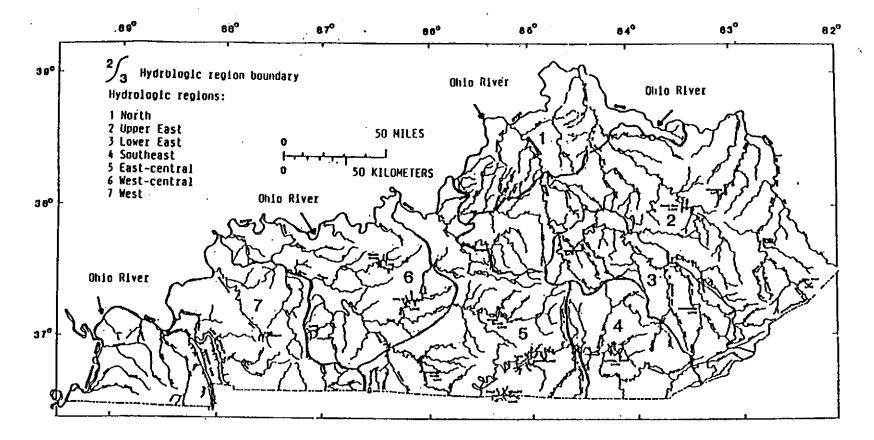
DELINEATION OF FLOOD REGIONS: USGS METHOD OF RESIDUALS

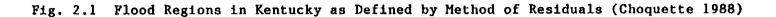
The U. S. Geological Survey currently employs the method of residuals to perform flood regionalization. The technique involves the use of residuals from a regression

equation relating a selected flow quantile (for example the 50-year flood level as obtained from an assumed probability distribution of the AMF series at each gauged site) to the physical and climatic characteristics of the watershed. The probability distribution employed is log-Pearson Type-III. This technique relies on the basic premise that the trends in the residuals reflect regional differences in the flood Thus, once a homogeneous region response of the watersheds. is delineated then the regression equation relating the flood response variable to the watershed characteristics will have residuals that can be attributed to pure chance. Unfortunately, the residuals contain both chance variation (time sampling error) and variation due to basin characteristics (model error) without a measure of the relative amounts of each (Riggs, 1973). This makes the delineation of homogeneous flood regions a difficult, if not an arbitrary, task to accomplish. Nonetheless, this procedure was used to delineate seven homogeneous flood regions for the State of Kentucky (refer to Figure 2-1) using flood data from all gauged sites used in the present study (i.e. both the data sets set aside for gauged and ungauged analysis). A regionalized skewness coefficient was used for estimating the 50-year flood quantile of the log-Pearson Type-III frequency curve fitted to annual peak flow data at each of the gauged sites.

VERIFICATION OF FLOOD REGIONS

a) Hydrological Characteristics of Flood Regions: For each of the clustering scheme and method of residuals the variation in important hydrologic characteristics (response and attribute variables) within and between regions are compared. Tables showing important statistics such as range, minimum and maximum, mean, and median are used for this purpose. These statistical characteristics of the





hydrological attributes at each site within a flood region provide a means to select clusters that may be similar or distinct from others. They will also indicate the type of watersheds that lie within each flood region.

b) Performance of Regionalized Flood Frequency Models: The accuracy of delineation flood regions can be further evaluated by examining the performance of the regionalized flood frequency model. Commonly used measures of performance are scaled values of bias and root mean squared error (RMSE). In this study, this is carried out using Monte Carlo simulation methods as oulined below.

- 1. For a selected probability distribution, estimate the at-site parameters using method of L-moments.
- 2. Generate normalized flood flow sequences, having the same record length in years as the historical systematic AMF record at the site, using a suitable a random number generator. A widely used random number generator referred to as RAND is used in this study. This is an IBM function that uses a multiplicative linear congruential method for generating a uniform set of pseudo-random numbers.
- 3. Using the IBM flood regionalization computer program (Hosking, 1988) a regionalized flood frequency model (for the selected probability distribution) is developed (for each region). A total of 100 simulation runs are made.
- 4. The regionalized flood frequency model developed in step 3 is used to estimate the flood quantiles at each site using the index-flood method described earlier.
- 5. The scaled bias in the estimate a flood quantile at each site is computed by taking the difference between the simulated value and the historical estimate (based on regionalized historical flood record) and dividing

this by the historical estimate. The RMSE uses the square of this scaled bias.

- 6. Steps 2 through 5 are repeated for each of the 100 simulation runs.
- 7. Based on the 100 simulation runs, regional average values of bias and RMSE are then computed using the corresponding estimates at each of the gauged sites within the region.

c) Discriminant Analysis: The success of any cluster analysis in identifying flood regions that are homogeneous within themselves but are distinct from the others depends to a large extent on the ability to discriminate between them. The variables to be used in discriminating between clusters or regions must be those that control flood response like the physical and climatic characteristics of the watershed(refer to nomenclature). Furthermore, the classification of an ungauged site (does not have observed AMF data) into a particular region can only be carried out using the attribute variables used in the discrimination process.

The power of discriminant analysis is measured by the correct reclassification of the gauged sites into their respective cluster regions that are originally identified in cluster analysis phase. A good discrimination can be obtained when the percentage misclassification of gauged sites is minimal. The success of accomplishing this objective depends on the attribute variables available for discrimination. The overall objectives of discriminant analysis in the context of flood regionalization are:

 a) To further explain the differences between cluster regions based upon hydrological variables (referred to as attribute variables) that affect and/or control flood response at each of the gauged site within a cluster region. This would further explain why the

flood regions (or clusters) are different with respect to the response variables used in the clustering process.

b) To use results from the discriminant analysis to classify ungauged sites that do not have their flood response variables defined.

d) Regression Analysis: The ultimate objective or purpose of regionalizing flood data is to develop regionalized relationships for predicting the flood response (at selected frequency levels) at both gauged and ungauged sites. For gauged sites, the regionalized relationship can be used together with at site information. The development of a regional equation for predicting flood response or quantiles within a given region can be accomplished using regression analysis by relating the flood level (dependent variable) with important hydrologic variables controlling flood response (independent or attribute variables). In the USGS method of residuals approach, this is accomplished by relating the log-Pearson Type-III flood quantile estimates at each gauged site within a region to hydrologic variables such as the geomorphic characteristics of the watershed. The regression analysis is carried out using log-transformed (base 10) data. The predictive capability of such equations is determined by examining the residual error expressed in percent (Tasker, 1978). Ideally, this error should be as low as possible.

COMPARISON OF FLOOD REGIONALIZATION METHODS

The main focus of this study, as stated earlier, is to compare the two methods of flood regionalization, namely, cluster analysis (Methods 1 and 2) and method of residuals (Method 3). In the following chapter homogeneous flood regions delineated under these two methods are compared with

those obtained by the method of residuals (refer to Figure 2.1) using the procedures discussed above. The following specific questions are addressed:

- a) How do the homogeneous flood regions delineated in the present study using cluster analysis, differ from those derived by the USGS Method of Residuals in terms of the watersheds and their hydrological characteristics?
- b) How well are the regions discriminated by the attribute variables under the two methods of regionalization?
 What are the most significant attribute variables that provide the maximum discrimination?
- c) For the selected probability distributions controlling flood response at each gauged site, how do the results of flood regionalization differ in terms of the performance of the regionalized flood frequency growth curves? What are the differences in the flood quantile estimates at each site? Flood quantile estimates from the log-Pearson Type-III distribution will also be included in this comparison.
- d) What are the differences in the regression equations that predict the flood quantiles (at various return periods) for each region using the two methods of regionalization? These regression equations are necessary for predicting flood quantiles at ungauged sites.

SPECIFIC RESEARCH PROCEDURES

Based on the overall procedures presented above, the following specific steps are followed in conducting this study:

 a) Hydrologic data, necessary for performing a regional flood frequency analysis, are obtained from the U.S.
 Geological Survey, Louisville District. These include observed annual flood data as measured at each of the the gauged sites (referred to as response variables) and physical, climatic and hydraulic characteristics of the watersheds that affect flood response (referred to as attributes).

- b) Probability distributions recommended for use in flood frequency analysis are selected after a careful review of previous research efforts. The following probability distributions, commonly employed in flood frequency analysis (Kuczera, 1982) are employed:
 - a) Generalized Extreme Value (GEV) and its special case, Extreme Value Type-I (EV1)
 - b) Wakeby

The parameters of the probability distribution will be estimated using the method of L-moments.

c) Cluster Analysis is then used to form homogeneous flood regions based upon important statistical properties of the normalized AMF series and the probability distribution selected in step (b). Properties such as the mean, standard deviation, coefficients of variation, skewness and kurtosis (L-moments) and the specific mean annual flood, QSP, and the parameters of the probability distribution, as estimated from L-moments, are used as indices to measure flood response of each watershed. The FASTCLUS procedure available in the Statistical Analysis System (SAS, 1985) is used to obtain clusters or groups. The purpose of this analysis is to place the gauged sites into groups or clusters such that gauges within a cluster have similar flood response and those in different clusters have dissimilar flood response.

- d) For the flood regions delineated in step (c) above, determine the most suitable regionalized probability distribution applicable to each of the gauged sites within the region using Monte Carlo simulation. This is based upon a performance criteria, such as the mean squared error and bias, that yield the most reliable estimates of extreme events. The simulation involves a detailed frequency analysis of the AMF series using regional parameters of the underlying probability distribution.
- e) For each flood region delineated in step (c) above, summarize and evaluate the trends in the hydrological characteristics and develop a regionalized flood frequency growth curves for a given probability distribution. Evaluate differences in the shapes of these growth curves between regions and relate this to differences in the hydrological characteristics.
- f) Perform Discriminant Analysis to distinguish between the clusters formed in step(c) based upon attribute variables such as the physical, climatic and other hydrologic characteristics of the watershed. The discriminant scores, associated with each of the attribute variables, are used to evaluate any misclassification of a gauged site into the homogeneous flood regions defined in step(c). This step will also identify the most important variables that affect or control flood response of a watershed and can later be used for developing flood prediction equations.
- g) Within each cluster, perform a stepwise regression analysis with using select flood quantile levels as the dependent variable and other watershed hydrologic attributes as the independent variables. This step will also identify the most significant attributes

variables controlling flood response of the watersheds within a cluster, and, additionally, provide a means to compare them with the set of attribute variables that contributed to the discriminant power between clusters as described in step (c) above. Compare the mean square and standard errors associated with the regression equations developed for each cluster region with similar equations obtained for the U.S.G.S. method of residuals flood regions. In this context, it must be emphasized that the actual gauges on each cluster will not be identical to those being used in the method of residuals study since the two methods are quite different in the manner in which the homogeneous flood regions are formed. However, the values of the errors associated with the regression equation within each cluster can be compared overall to those obtained from the method of residuals in order to determine the most suitable method of regionalization.

CHAPTER 3

DATA AND RESULTS

DATA ACQUISITION

Annual maximum floodpeak data (AMF series) was retrieved from WATSTORE by the U.S. Geological Survey, Louisville District office. Additional hydrologic data pertaining to each watershed corresponding to the gauged streamflow sites was provided by the U.S. Geological Survey office in Louisville. This data constitutes a part of the information on the attribute variables (or independent variables) to be used in the regionalization study. Additional geomorphic variables for each watershed may be necessary to further improve the regionalization process. Such data was not readily available at the completion of this report.

The following is a detailed list of hydrologic, physical and meteorologic data that is used in the flood regionalization study.

- The systematic historic AMF record at each of the gauges in the State of Kentucky. Only gauges located in watersheds with drainage areas less than 1000 square miles and having at least 7 years of flood data is used in the analysis.
- 2) Physical characteristics affecting or controlling the flood response of the watershed in which the gauge is located. This includes watershed contributing drainage area, A_c, length, B₁, shape index, B_s, average slope, B_s, elevation, soil type, and land use (percent impervious area etc.),

and main channel length, L_c , sinuousity, S_s , and slope, S_c . Geomorphic data such as the number and average length of streams of different orders (for computing geomorphic properties of each watershed such as stream order, stream frequency, drainage density, form factor and bifurcation ratio), and the time of concentration were not readily available at the completion of this study.

3) Climatic data such as seasonal (dry and wet periods) and type of rainfall characteristics experienced in each of the watersheds. The only variable available at the time of this study was the mean annual rainfall.

The list of flood response variables (dependent variables) and the watershed attribute variables (independent variables), to be used in the regionalization study is shown at the end of this report under nomenclature. Pertinent statistical of data corresponding to these variables, as defined at each of the 253 gaging sites, is included in Table A.1, Appendix A. The values of the response variables are derived by computing important statistics of the normalized AMF data for each gauged site. These statistics, either individually or in combination, will be used in defining homogeneous flood regions using cluster analysis as presented in the following sections.

DELINEATION OF CLUSTER FLOOD REGIONS

Using FASTCLUS algorithm, a detailed cluster analysis is carried out using the response variables outlined in the previous section and in Chapter 2 with the following objectives.

1) To obtain optimum number of clusters or regions

that are physically realistic for representing flood experience for the State of Kentucky.

- 2) The number of clusters selected must satisfy at least one of the several available cutoff criteria. This would ensure that each cluster is homogeneous within itself but heterogeneous with respect to other clusters.
- 3) The number of gauged sites within a cluster must be sufficiently high in order to permit any statistical analysis.
- 4) The clusters must lend themselves to maximum possible discrimination based on the attribute variables (hydrological characteristics other than those based on AMF data). This would maintain the hydrologic distinction between the cluster regions.
- 5) The misclassification of the gauged sites already grouped and the ungauged sites to be assigned to a cluster region must be minimal.

With the above objectives in mind, results from cluster analysis using FASTCLUS algorithm are initially screened for the most suitable response variables to be used for further analysis. These results suggest that independent clusters or flood regions can be successfully formed using the statistical L-moments, LCV, LSK, LKUR of the normalized annual peak flow data, the parameters of the selected probability distribution and the specific mean annual flood, QSP taken individually or in combination. Clustering on physical characteristics of the watershed gave cluster regions that could not be discriminated well based on the flood response variables.

As expected, the composition of each cluster and the optimum number of clusters that can be extracted and discriminated (based upon attribute variables associated with the watersheds in which each of the gauged sites is

located) continues to depend heavily on the type and number of response variables used in the analysis. Consequently, the final choice of clustering schemes, incorporating different response or clustering variables, is based upon the overall performance of each flood region. The following sections discuss results of all the clustering schemes and techniques used to delineate and evaluate the flood regions.

a) Clustering Cases: Twelve clustering schemes are adopted initially for further examination. Table 3.1 summarizes the results of the FASTCLUS clustering procedure for the various clustering schemes. Case 13 shown in this table applies to USGS regions, as delineated using method of residuals, and is included for the purpose of comparing the two method of regionalization. These twelve cases, as shown in Table 3.1, involve clustering with the response variables L-moments, namely, coefficients of variation, LCV, skewness, LSK, and kurtosis, LKUR, respectively, the specific mean annual flood, QSP, and the parameters of the EV1 (MEVL and AEVL), GEV (MGVL, AGVL and KGVL) and Wakeby (MWKL, AWKL, BWKL, CWKL, and DWKL) distributions. Each case is included in the study with a specific purpose. For example, for the clustering cases involving L-moments (Cases 1-3), Case 1, with clustering variable, LCV, would be appropriate for 2-parameter flood frequency models that require location and scale parameters to characterize the model completely. It must be emphasized, that the use of normalized AMF data standardizes the first moment (mean), characterizing the location, to 1.0. Since the coefficient of variation, LCV, reflects the dispersion (or scale) effects present in the flood data, this statistic would be totally adequate to describe a flood frequency model involving location and scale parameters. For example, the EV1 distribution used in this study can be characterized completely by LCV. In contrast, a five parameter flood frequency model like the Wakeby would require all L-moments, LCV, LSK and LKUR and

No.	Cluster Variables (No. of Clusters	R ²	ccc	No. of Sites in Each Cluster Region
1	LCV	6	0.953	-6.66	78,33,42,20,70,10
2	LCV, LSKEW	6	0.830	-1.17	66,45,16,57,31,38
3	LCV, LSKEW, LKUR	6	0.766	6.36	38,49,19,43,66,38
4*	LCV, QSP	5	0.759	-4.50	89 16,93,30,25
5*	LCV, LSKEW, QSP	5	0.689	1.97	79,17,75,44,38
6	LCV, LSKEW, LKUR,QSP	• 5	0.611	4.82	26,26,73,88,40
7	MEVL, AEVL	6 ·	0.953	23.80	79,34,41,20,70,9
8	MGVL, AGVL KGVL	5	0.705	3.65	74,30,29,46,74
9	MWKL, AWKL, BWF CWKL, DWKL	CL 2	0.215	7.41	44,209
10*	MEVL, AEVL, QSP	5	0.775	12.24	43,10,91,79,30
11*	MGVL, AGVL, KGVL, QSP	· 6 _	0.646	4.27	81,21,40,15,68,28
12 }	WKL, AWKL, BWKI CWKL, DWKL, QSI		0.287	6.57	5,12,236
13#	USGS REGIONS	7		- :	32,68,26,20,38,31,38

- .

TABLE 3.1. Clustering Characteristics of Cases Examined in the Study

one higher order moment, LBMD. Cases 7-9 correspond to Cases 1-3 with the exception that the actual at-site parameters (as estimated from L-moments) of the appropriate flood frequency model are used as clustering variables. Hence, Case 1 would correspond to Case 7 since the estimation of EV1 parameters require LCV (for normalized AMF flows). Cases 4-6 and 10-12 are similar to the above cases but include an important clustering variable, namely the specific mean annual flood, QSP. Unlike all the other clustering variables, which describe the underlying flood frequency distribution, the specific mean annual flood describes the flood potential of each watershed. An examination of at-site estimates of QSP for the 253 gauged sites in Kentucky indicates that its value decreases as the size of watershed increases.

The relative performance of the above 12 clustering cases is evaluated in detail using the following results.

- 1. Results of the cutoff criteria for choosing optimum number of clusters,
- 2. Trends in the hydrological characteristics and regionalized frequency growth curves,
- 3. Performance of the regional flood frequency model using simulation,
- 4. Results of discriminant analysis, and
- 5. Results of regression analysis relating flood quantiles to watershed physical and climatic characteristics.

b) Selection of number of cluster regions and cases: Since one of the main objectives of cluster analysis, in the context of flood regionalization, is to delineate homogeneous flood regions that can be distinguished from each other, the number of clusters obtained must not be too few or large. With this in mind, several cutoff criterion or stopping rules, as discussed in Chapter 2, are used to determine the optimum number of cluster regions. An

application of the stopping rules to the 12 clustering schemes gave results shown in Table 3.1. For all schemes the CCC criterion showed a peak or trough value going from larger to smaller number of clusters than the optimum number of clusters (refer to column 6 of Table 3.1) and the R^2 was quite high indicating a clear choice of the optimum number of clusters. The inclusion of QSP as a clustering variable changed the optimum number of cluster regions from 6 to 5 with the exception of the case when the GEV parameters are used in the clustering. Clustering on Wakeby parameters gave only 2-3 regions and gave the worst overall performance compared to all clustering cases examined in this study. Hence, the Wakeby probability distribution is not considered suitable for regionalizing flood data for the State of Kentucky and is dropped from further consideration. Amongst the remaining schemes, the inclusion of QSP as a clustering variable (refer to scheme numbers 4, 5, 10 and 11 in Table 3.1) improved, although marginally, the overall performance. Consequently, all results discussed in the following sections pertain to the following four cases (marked by an asterisk "*" in Table 3.1) that are finally selected from the twelve clustering schemes. These cases incorporate the flood regionalizations Methods 1 and 2.

Case 1 :	Clustering with LCV and QSP (Method 1)
Case 2 :	Clustering with the Extreme Value Type-1
	probability distribution parameters (MEVL and
	AEVL) and QSP (Method 2)
Case 3 :	Clustering with LCV, LSK and QSP (Method 1)
Case 4 :	Clustering with the Generalized Extreme Value
	parameters (MGVL, AGVL and KGVL) and QSP
	(Method 2)

As mentioned in the previous section, Case 1 and Case 2 are similar since the clustering variable LCV is adequate to estimate the parameters MEVL and AEVL of the EV1

distribution. Also, Case 3 and Case 4 are similar since the clustering variables LCV and LSK are used to estimate the parameters MGVL, AGVL and KGVL of the GEV distribution. All the four clustering cases gave, by and large, disjoint cluster regions as illustrated in the bi-variate plots shown The numbers shown on these figures are in Figures 3.1-3.13. cluster numbers. It is obvious from these figures that the the overlap between cluster regions increases as the number of clustering variables increase (refer to Case 4, Figures 3.8-3.13). The bi-variate plot of EV1 parameters, MEVL versus AEVL (refer to Fig. 3.3), shows an inverse linear relationship suggesting an increase in the location parameter (mode) as the scale parameter decreases. The bi-variate plot involving L-moments, as in Case 3, illustrates that LCV is directly proportional to LSK (refer to Fig. 3.6).

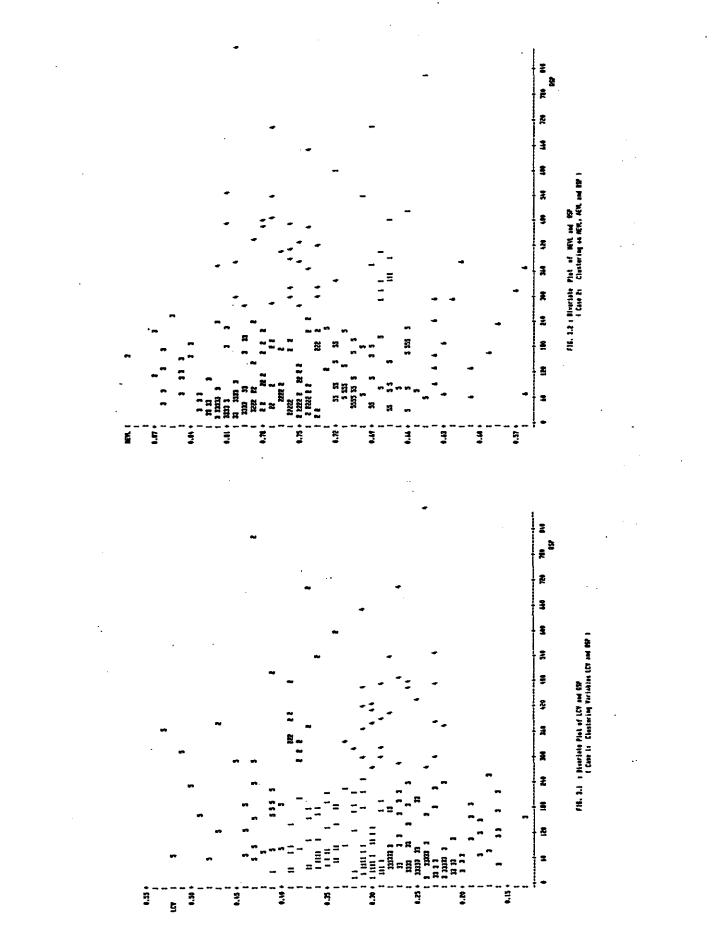
The total number of gauged sites classified into each of the cluster regions for the above four cases is shown in the last column of Table 3.1. The smallest number actual sites within a cluster is 10 (Case 2) which is adequate for performing any statistical analysis within the region.

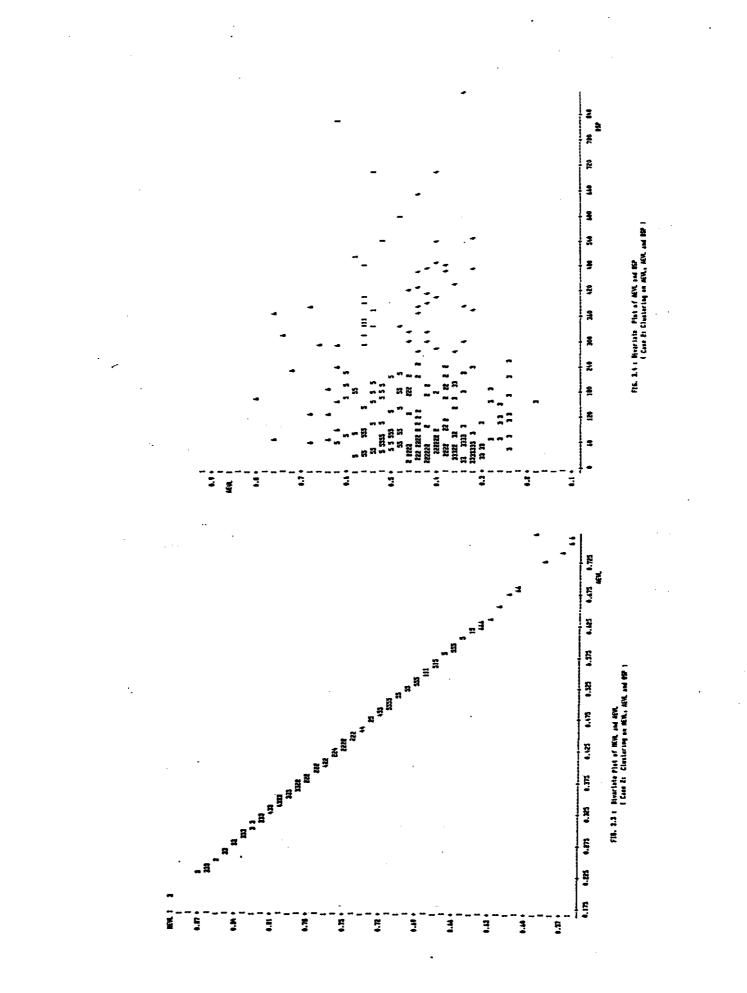
The number of gauged sites (not the actual gauges) assigned to a particular cluster depends on the clustering variables used in the analysis. This is illustrated in Tables 3.2-3.7. Using Cases 1-4 (clustering with response variables LCV, LSK, EV1 and GEV parameters and QSP), these tables show the number of gauges reassigned when the clustering case is changed to one of the remaining cases. Each row reflects the number of gauged sites reassigned to the cluster numbers shown in the columns when clustering is carried out using any other case in lieu of the one shown on left hand side. For instance, the first row in Table 3.4 shows that of the 89 gauged sites (refer to last column of Table 3.4) assigned to cluster 1 when using clustering variables, LCV and QSP (Case 1), 65 sites are reassigned to Cluster 1 when using LCV, LSK and QSP as clustering

variables (i.e. Case 3), 21 gauged sites are reassigned to Cluster 3, and 3 sites to Cluster 4. Thus, there is a clear evidence of movement in the gauged sites between clusters when Case 1 and 3 are compared against each other. Α similar comparison of Case 1 (clustering with LCV and QSP) versus Case 2 (clustering with EV1 parameters and QSP) and Case 3 (clustering with LCV, LSK and QSP) versus Case 4 (GEV parameters and QSP) respectively (refer to Tables 3.2 and 3.5), also suggests movement, although to a lesser degree, between cluster regions. Thus, the cluster regions delineated using the L-moments or parameters tend to be dependent on the type and number of clustering variables used. The effect of using different clustering variables (although standardized) on the hydrological composition of cluster regions delineated is illustrated further in the following sections.

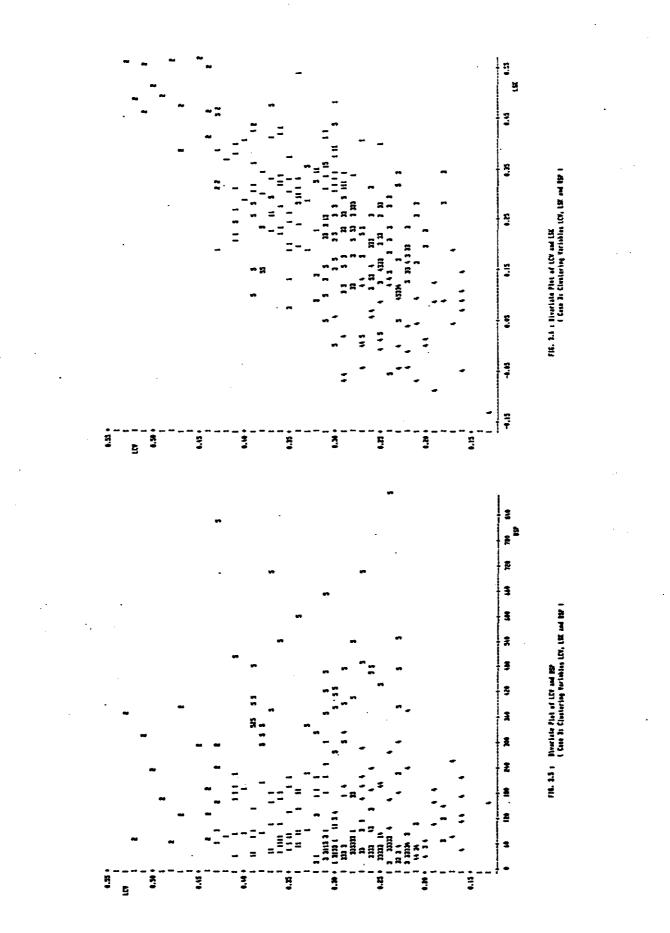
c) Comparison of Cluster and USGS Regions: The seven flood regions delineated by the USGS using the method of residuals (refer to Figure 2.1), are quite different in terms of the · actual gaged sites when compared to those obtained by cluster analysis. Since cluster regions are not coincident with any geographic or hydrologic boundaries, they can not be illustrated in a convenient manner like the USGS regions of Figure 2-1. Furthermore, the total and the individual gauged sites incorporated within a region vary considerably. This is clearly evident from a comparison of the USGS method of residuals regions with the cluster regions obtained under each of the four cluster schemes (cases 1-4). For example, Tables 3.8-3.11 compares the USGS regions with those obtained under clustering Cases 1-4. In these tables, the rows represent the cluster regions for a particular case with the total gauged sites within each region shown in the last column. In the same manner, the columns represent each of the seven USGS regions (as delineated using method of residuals) with the total gauged sites within a cluster

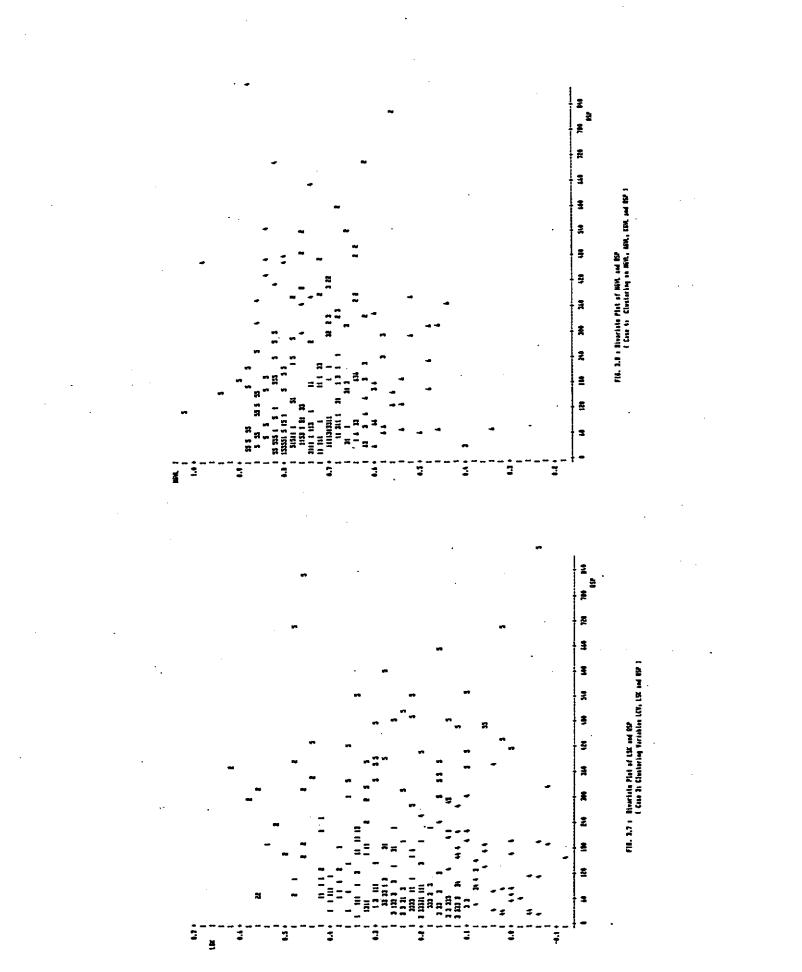
shown in the last row. An examination of these tables indicates, as expected, significant movement of gages between the cluster and USGS regions.

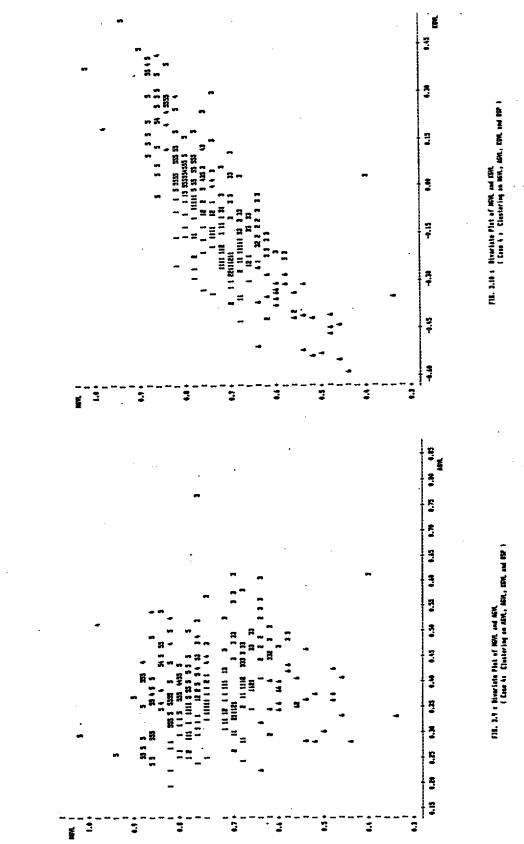


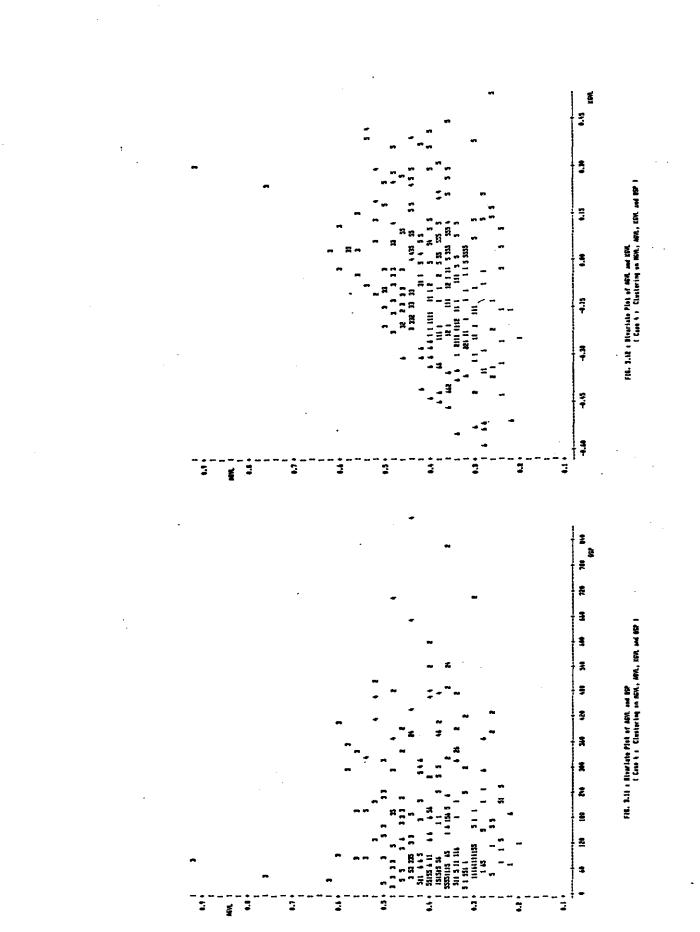


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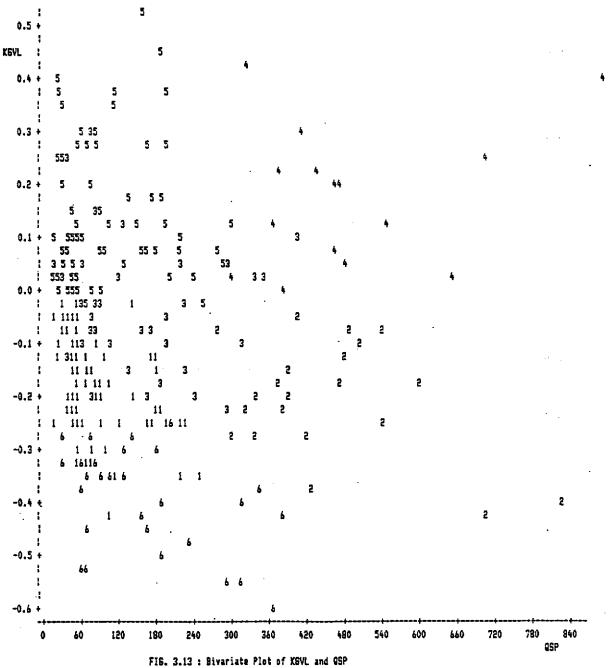








.



(Case 4 : Clustering on MGVL, AGVL, KGVL QSP)

					ME	л.	VEAT	£ 0	SP				
			1		2		3		4	_ _	5		Total
	1	I	11	l	0	1	78	I	0	I	0	I	89
LCV	2	1	7	I	9	1	0	I	0		0		16
6	3	I	0	1	0		13	Ι	79	!	1	I	93
QSP	4	I	0	1	1	1	0	` I	0	1	29	I	30
	5	1	25	l	0	l	0	·	0		0	1	25
To	tal	1	43	1	10		91	l	79	1	30	1	253

TABLE 3.2	l. Compa	rison d	of Act	ual Number	of	Gauged	Sites
Assigned	Between	Case 1	and 2	Clustering	r Se	chemes	

					G	EV	Para		ater	8 2	and Q	SI	?		
•	<u></u>	_ _			2		3		4	_ _	5		6		Total
	1	1	47	I	0	1	26	I	0	ł	7	1	9	I	89
LCV	2	1	0	1	9	1	5	Ι	0	1	0	I	2	I	16
2	3	1	34	I	0	ł	1	1	0	1	58	1	0	1	93
Q5P	4		0	I	12	I	0	1	15	1	3	I	0	I	30
	5	I.	0	I	0	1	8	1	0	I	0	I	17	I	25
To	tal	1	81	I	21	t	40	1	15	1	68	1	28	1	253

TABLE 3.3. Comparison of Actual Number of Gauged Sites Assigned Between Case 1 and 4 Clustering Schemes

TABLE 3.4. Compariso

	Comparison o				
Assigned Be	stveen Case 1	and 3 Clu	ustering :	Schenes	

					LC	V, 1	SKEW	£ ()SP				
			1	!	2		3		4		5	1	Total
	1	1	65	1	0	1	21	1	3	1	0		89
LCV	2	I	0	1	2		0	I	0	1	14	I	16
£	3	Ι	3	1	0	1	54	1	36		٥	1	93
SP	4	ŀ	1	1	0		0	1	5	1	24		30
	5	I	10	1	15		0		0		0	.	25
Ta	tal	1	79		17		75	1	44		38	·	253

TABLE 3.5	5. Compa	rison (of Actu	al Number	of	Gauged	Sites
Assigned	Batween	Case 3	and 4	Clustering	y Se	chemes	

.

					GE	VF	ara	Pet	ers:	ar	nd Q	SP			
		_ _	1		2		3	_ _	4		5		6		Total
	1	1	39	I	1	 .	26	I	0	I	0	I	13	L	79
LCV, LSKE	2	1	0	1	0	I	2	İ	0	I	0		15	1	17
£,	3	I	42	I	0	1	7	I	0	I	26	1	0	I	75
Q\$₽	4	1	0	I	0	l	0	I	2	I	42	1	0	I	44
	5	I	0	I	20	I	5	1	13	1	0	1	0.	I	38
то	tal	1	81	Ι	21	I	40	1	15	1	68	1	28	1	253

				Ev	T ta	, a se	ters	ano	i yar				
			1		2		3		4		5		Tota:
V	1	1	0	1	0	I	54	1	27	1	0	1	8
	2	I	1	I	8	۱	0	I	0	I	12	I	2
	3	I	17	1	1	I	22	I	0	I	0	I	4
	4	I	0	Ĩ	1	l	0	I	0	I	14	1	1
	5	1	0	I	0	ł	12	I	52	I	4	۱	6
P	6	I	25	1	0	I	3	I	0	I	0	ł	2
To	tal	1	43		10	I	91	1	79	1	30	1	25

TABLE 3.6. Comparison of Actual Number of Gauged Sites Assigned Between Cases 4 and 2 Clustering Schemes .

MARTER 1	0 <i>C</i> om	and som of	1 Lotus 1	Number of	of Gauged	Olton .
TAOLE J.	o, comp	WCTROU OT	, AGLUAL	NUBLOBL C	or Gaugeo	STICS
3			Beaters	10 11) and USGS	* Baadaaa
veer duer	Dercassi	CIUSCOF.	redrous	(Case I)) 64144 4363	xediour

							σ.	. 5.	. (3. 8	3.	Reg	ji	ons				
				1		2		Э		4		5		6		7	_	Total
		1	1	5	1	29	I	14	1	6	l	15	I	5	I	15	I	89
cv		2	Ι	3	I	0	l	2	l	0	1	3	I	3	1	3	1	16
2		3	1	14	1	28	I	6	I	12	1	11	ł	15	I	7	I	93
SP		4	l	6	1	4	1	1	1	0	1	3	1	6	1	10	1	30
	•	5	I	4	I	7	I	3	١	2	I	6	I	2	I	1	1	25
	Tot	al	1	32	I	68	1	26	1	20	1	38	I	31	1	38		253

TABLE 3.7. Comparison of Actual Number of Gauged Sites Assigned Between Cases 3 and 2 Clustering Schemes
PVI Parameters and OSP

				EV	I Pa	e 7. b 1 (6	Cers	ano	l QSP				
		_ _	1		2		3		4		5		Total
	1	I	21	1	0	1	55	1	2		1	1	79
LCV, LSK	2	I	17	I	0	I	0	l	0	1	0	1	17
£	3	I	0	I	0	Ι.	32	ł	43	1	0	I	75
QSP	4	I	0	1	0	1	4	I	34	I	6	1	44
	5	1	5	1	10	1	0	I	0	!	23	1	38
Tot	al	I	43	l	10	1	91	I	79	1	30	1	253

TABLE 3.9. Comparison of Actual Number of Gauged Sites Assigned Between Cluster Regions (Cases 3) and USGS Regions

						τ	J. 1	5.	G.	S	. R	eg:	lons	5			
			1	_ _	2	1_	3		4		5		6		7		Total
	1	1	7	l	21	I	13	ł	6	ł	14	I	5	I	13	I	79
H	2	I	4	I	4	I	1	I	0	1	6	I	1	1	1	1	17
	3	1	5	I	29	I	8	1	9	t	9	I	9	ļ	6	I	75
	4	1	7	1	13	1	2	1	5	I	5	I	8	١	4	I	.44
	5	I	9		1	I	2	1	0	1	4	1	8	1	14	Ţ	38
To	tal		32		-68	÷Ê-	-26	1	20	 	38		31	!	38		253

						۱	σ. :	5.	G.	S	. Re	eg:	ions	5			
	<u> </u>		1		2		3	1	4		5		6		7 	1	Total
ev1 P	1	1	8	١	9	ł	6	1	2		12	1	3	1	3	ł	43
A R A	2	ļ	1	1	0	I	0	1	0	l	0	l	3	1	6	1	10
M E T	3		4	1	32	I	14	1	9	!	14	I	4	. 1	14	1	91
e R S	4	ļ	12		23	1	5		9	1	9	1	15	1	6		79
æ	5	-	7	I	4	1	1	I	0	1	3	!	6		. 9	l	30
QSP	Total	1	32	1	68		26		20		38	1	31	1	38	!	253

TABLE 3.10. Comparison of Actual Number of Gauged Sites Assigned Between Cluster Regions (Cases 2) and USGS Regions

TABLE 3.11.	Comparison o	f Actual	Number o	f Gauged	Sites
Assigned Bet	ween Cluster	Regions	(Cases 4)	and USG	S Regions

						1	σ. :	5.	G.	S.	. Re	eg:	ions	3		•	
	·	_ _	1		2		3		· 4		5	_[_	6		7	_ _	Total
GEV	1	۱	6	ł	26	ł	10	I	9	1	10	!	7	I	13	1	81
Р А R	2	ļ	5	ľ	2	Į	0	I	0	1	2	1	4	1	8	}	21
a M E	3		4	1	11	1	7	1	3	I	9	1	2	1	4	1	40
T E R	4	ł	2	l	1		1	1	0	1	2	I	3	i	6	l	15
s £	5		10	l	21	1	4	1	8	1	8	1	13	1	4	1	68
QSP	6	-	5	1	7	ł	4	1	0	1	7	1	2	I	3	ļ	28
То	tal	-	32		68		26	!	20	1	38	l	31	1	38	I	253

DEVELOPMENT OF FLOOD FREQUENCY GROWTH CURVES

The procedure for developing a regionalized flood frequency growth curve was presented in Chapter 2. For the four clustering cases (Case 1-4), a separate regionalized flood frequency growth curve is developed for the EV1 and GEV probability distributions using historical systematic annual maximum floodpeak series (AMF series) from each of the gauged sites within a cluster region. The index-flood procedure presented in Chapter 2 is applied to accomplish the regionalization. The regionalized weighted (by the record length at each site) average L-moments and the corresponding EV1 and GEV parameters are shown in Tables 3.12 and 3.13 for each of the cluster regions delineated under the four clustering schemes. Similar data for the USGS regions are included for comparative purposes. For the USGS regions, the regionalized EV1 and GEV distributions are fitted using the method of L-moments. The actual gauged sites within each of the seven regions are identical to those contained in the regions delineated by the method of residuals (Choquette, 1988).

The EV1 and GEV regionalized flood frequency growth curves developed from the parameters in Table 3.13, are illustrated in Figures 3.14-3.21. It is important to note that the cluster numbers assigned to each region will change from case to case. Thus, cluster region number 5 in Figure 3.14 for Case 1 is not the same as cluster number 5 in Case 2. These numbers are arbitrarily assigned during the clustering process. Similar curves are developed for the USGS Method of Residuals regions are shown in Figures 3.22-3.23 for the EV1 and GEV distributions, respectively. The vertical scale (showing normalized discharge values) for the EV1 (Gumbel) distribution is drawn to half the scale than the one used for the GEV in order to improve the clarity of these frequency growth curves.

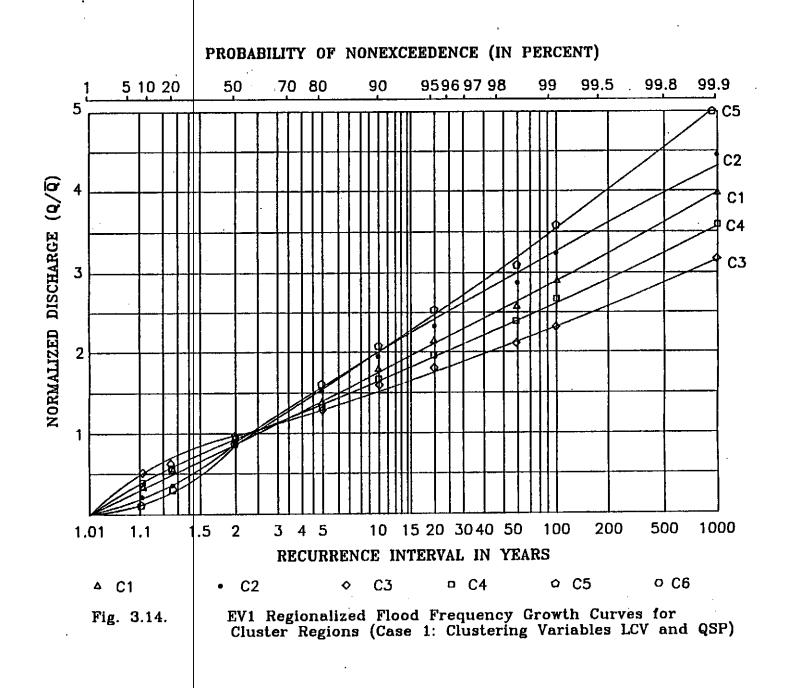
TABLE 3.12. Comparison of Regional Average L-Moments Estimated Using Normalized Historic AMP Data TABLE 3.13. Comparison of Regional Average Parameters of EV1 and GEV Probability Distributions Fitted to Normalized Historic AMF Data /

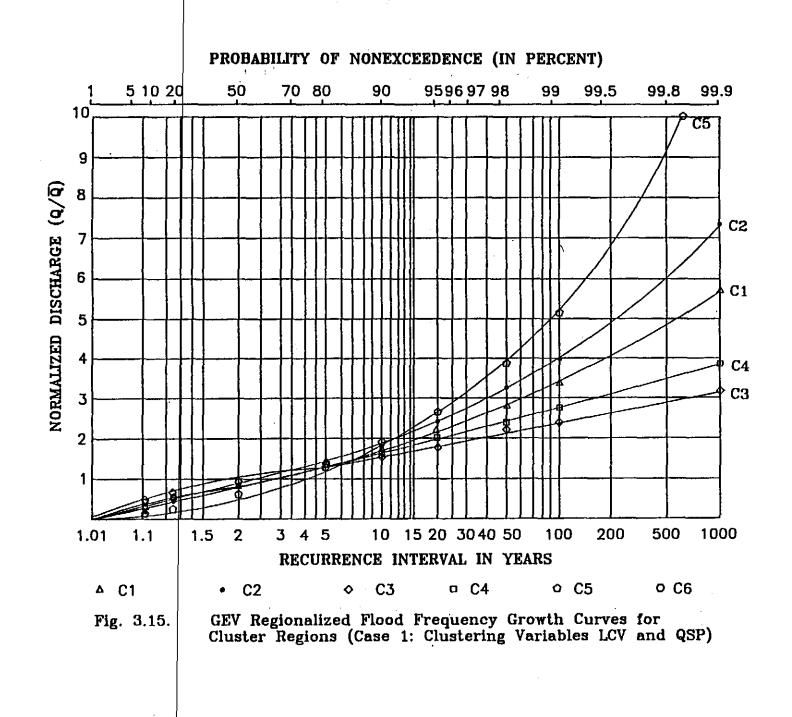
1.0000 1.0000 1.0000 1.0000	0.2383 0.2823 0.3242 0.3862 0.4432	Ж(I.SK) 0.1760 0.1908 0.2758 0.3050 0.4035 • (MEVL, XEVL	N(LKUR) 0.1810 0.1839 0.1914 0.1900 0.2784	H (LBHD) 0.0781 0.0786 0.1016 0.0809	Region No. * Cluster Reg Case 1: Clu J 4 1		AEVL 	HGVL. 0.80 0.76 0.70	AGVL 0.34 0.39	KGVL -0.01 -0.94
ng with LCM 1.0000 1.0000 1.0000 1.0000 1.0000 1.0000 1.0000 1.0000 1.0000 1.0000	<pre># and QSP 0.2383 0.2823 0.3242 0.3862 0.4432 0.4432 0.3862</pre>	0.1760 0.1988 0.2758 0.3058 0.4035	0.1810 0.1839 0.1914 0.1900 0.2764	0.0781 0.0786 0.1016 0.0809	Case 1: Clu J 4 1	stering on) 0,80 0,76	0.34 0.41	0.76	0.39	
ng with LCV 1.0000 1.0000 1.0000 1.0000 1.0000 1.0000 1.0000 1.0000 1.0000 1.0000	0.2383 0.2823 0.3242 0.3862 0.4432 parameter	0.1988 0.2758 0.3058 0.4035	0.1839 0.1914 0.1900 0.2784	0.0786 0.1016 0.0809	3 4 1	0,80 0,76	0.34 0.41	0.76	0.39	
1.0000 1.0000 1.0000 1.0000 1.0000 1.0000 1.0000 1.0000 1.0000	0.2383 0.2823 0.3242 0.3862 0.4432 parameter	0.1988 0.2758 0.3058 0.4035	0.1839 0.1914 0.1900 0.2784	0.0786 0.1016 0.0809	4	0.76	0.41	0.76	0.39	
1.0000 1.0000 1.0000 1.0000 1.0000 1.0000 1.0000 1.0000	0.2823 0.3242 0.3862 0.4432 parameter	0.1988 0.2758 0.3058 0.4035	0.1839 0.1914 0.1900 0.2784	0.0786 0.1016 0.0809	1	•				+0.04
1.0000 1.0000 1.0000 ng on EV1 p 1.0000 1.0000 1.0000	0.2823 0.3242 0.3862 0.4432 parameter	0.1988 0.2758 0.3058 0.4035	0.1839 0.1914 0.1900 0.2784	0.0786 0.1016 0.0809		0.73	0 47	0.70		
1.0000 1.0000 1.0000 ng on EV1 p 1.0000 1.0000 1.0000	0.3242 0.3862 0.4432 Parameter	0.2758 0.3058 0.4035	0.1914 0.1900 0.2784	0.1016 0.0809	• •	v	0.4/	V4/V	0.40	-0.16
1.0000 1.0000 ng on EV1 g 1.0000 1.0000 1.0000	0.3862 0.4432 Parameter	0.3058 0.4035	0.1900 0.2784	0.0809	4	0.68	0.56	0.63	0.44	-0.20
1.0000 ng on EVI g 1.0000 1.0000 1.0000 1.0000	0.4432 Parameter	0.4035	0.2784		5	0.63	0.64	0.55	0.42	-0.33
ng on EV1 1.0000 1.0000 1.0000 1.0000	parameter			0.1657	-					
1.0000 1.0000 1.0000 1.0000	J	MEVL, AEVL		0.1037	Case 2: Clu	stering on 1	EV1 parameter	S (MEVL, AEVL)	and QSP	•
1.0000 1.0000 1.0000	0.2310) and QSP			-	-			
1.0000 1.0000 1.0000	0.2310				4	0.81	0.33	0.81	0.34	0.01
1.0000		0.1641	0.1829	0.0741	5	0.77	0.40	0.76	0.39	-0.05
1.0000	0.2801	0.2008	0.1834	0.0743	3	0.74	0.45	0.71	0,39	-0.14
	0.3116	0.2637	0.1869	0.1001	2	0.70	0.52	0.66	0.43	-0.17
1.0000	0.3621	0.2837	0.2074	0.1097	1	0.65	0.61	0.58	0.43	-0.29
	0.4196	0.3703	0.2451							
	1				Case J: Clu	stering on 1	LCV, LSK and	QSP		
ng with LCV	/, ISK and	QSP								
]				4					0.19
	0.2229	0.0530	0.1324	0.0519	-					-0.05
1.0000	0.2613	0,2005	0.1713	0,0784	-					-0.11
1.0000	0.3224	0.2420	0.1853		1	0.72	0.49	0.68	0.38	-0.22
1.0000	0.3383	0.3187			2	0.62	0,67	0.52	0.38	-0.42
1.0000	0.4615	0.4672	0.3275							
					Case 4: Clu	stering on (SEV parameter	s (HGVL, AGVL,	, KGVL) A	nd QSP
ng on GEV p	arameters	i (MGVL, AGVL	, KGVL) and	QSP	_		• • •			
				•	Ş					0.10
			0.1379	0.0527	4					0.15
				0.0644	-					-0.18
				0,1145						-0.22
		0.3218	0.2319		-					-0.10
		0.2335	0.1385	0.0700	6 _i	Q.65	0.61	U. 50	0.30	-0.40
1.0000	0.4233	0,4544	0.3236.	0.1868	nece					
					napa kediou	5 8				
	J				6	0.78	0.39	0.76	0.36	-0.08
1 0000	0.000				1	-				-0.15
					1					-0.05
					1					-0.09
					4					-0.09
										-0.15
					2					-0.15
					2	1.1	V.40	v./u	U.JQ	- U + 4 /
1.0000	0.3185	0.2852	0.2061	0.1042	J MEUR MA	UT. = loostl	n parameters	/modest + 180	. NOVI -	scale
					MEVL, MG	AP = TOCULIO	·	(HOUES); ACVL	1 YOAP -	JUAIN
ed in incr	easing st	sepness of th	e correspoi	nding flood	paramete	Laj ang Kuvi	- suepe par	taenness of th		nonding f
th Curves	(1.e. inc	reasing LCV o	or LSK)		* Regions	arrangea in	increasing B	Laepness of th	IC COLLEN	hougened r
	1.0000 ng with LCV 1.0000 1	1.0000 0.4196 ng with LCV, LSK and 1.0000 0.2229 1.0000 0.2613 1.0000 0.3224 1.0000 0.3224 1.0000 0.4615 ng on GEV parameters 1.0000 0.2735 1.0000 0.2735 1.0000 0.2735 1.0000 0.3566 1.0000 0.2698 1.0000 0.2698 1.0000 0.2698 1.0000 0.2698 1.0000 0.2781 1.0000 0.2839 1.0000 0.2839 1.0000 0.3115 1.0000 0.3185 .0000 0.3185	1.0000 0.4196 0.3703 mg with LCV, LSK and QSP 1.0000 0.2229 0.0530 1.0000 0.2613 0.2005 1.0000 0.3224 0.2420 1.0000 0.3383 0.3187 1.0000 0.4615 0.4672 mg on GEV parameters (MGVL, AGVL, 1.0000 0.2379 0.1054 1.0000 0.2379 0.1054 1.0000 0.2827 0.2877 1.0000 0.3130 0.3218 1.0000 0.3566 0.2335 1.0000 0.4623 0.4544 1.0000 0.2698 0.2230 1.0000 0.2839 0.2265 1.0000 0.2839 0.2265 1.0000 0.3115 0.2695 1.0000 0.3115 0.2695 1.0000 0.3185 0.2852 .0000 0.3185 0.2852	1.0000 0.4196 0.3703 0.2451 ng with LCV, LSK and QSP 1.0000 0.2229 0.0530 0.1324 1.0000 0.2613 0.2005 0.1713 1.0000 0.3224 0.2420 0.1853 1.0000 0.3383 0.3187 0.2178 1.0000 0.4615 0.4672 0.3275 ng on GEV parameters (MGVL, AGVL, KGVL) and 1.0000 0.2373 0.1054 0.1379 1.0000 0.2373 0.0767 0.1239 1.0000 0.2827 0.2877 0.2218 1.0000 0.3383 0.3118 0.2319 1.0000 0.3566 0.2335 0.1185 1.0000 0.4231 0.4544 0.3236 1.0000 0.2698 0.2230 0.1867 1.0000 0.2698 0.2230 0.1867 1.0000 0.2698 0.2230 0.1867 1.0000 0.2819 0.2037 0.1582 1.0000 0.2839 0.2265 0.1994 1.0000 0.2839 0.2265 0.1994 1.0000 0.3115 0.2852 0.2061	1.0000 0.4196 0.3703 0.2451 0.1359 ng with LCV, LSK and QSP 1.0000 0.2229 0.0530 0.1324 0.0519 1.0000 0.2613 0.2005 0.1713 0.0784 1.0000 0.3224 0.2420 0.1853 0.0809 1.0000 0.3187 0.2178 0.1176 1.0000 0.4615 0.4672 0.3275 0.1847 ng on GEV parameters (MGVL, AGVL, KGVL) and QSP 1.0000 0.2735 0.1054 0.1379 0.0527 1.0000 0.2827 0.2877 0.2218 0.1145 1.0000 0.3130 0.3218 0.2319 0.0952 1.0000 0.3566 0.2335 0.1385 0.0700 1.0000 0.4621 0.4544 0.3236 0.1868 1.0000 0.2781 0.2728 0.2139 0.0989 1.0000 0.2810 0.2037 0.1582 0.0801 1.0000 0.2839 0.2265 0.1994 0.1044 1.0000 0.2839 0.2265 0.1994 0.1044 1.0000 0.3115 0.2695 0.1810 0.0887 1.0000 0.3185 0.2852 0.2061 0.1042 med in increasing steepness of the corresponding flood	1.0000 0.4196 0.3703 0.2451 0.1359 ng with LCV, LSK and QSP 4 1.0000 0.2229 0.0530 0.1324 0.0519 3 1.0000 0.2229 0.0530 0.1324 0.0519 3 1.0000 0.2224 0.2420 0.1853 0.0809 1 1.0000 0.3224 0.2420 0.1853 0.0809 1 1.0000 0.324 0.2470 0.1853 0.0809 1 1.0000 0.34615 0.4672 0.3275 0.1847 Case 4: Clux ng on GEV parameters (MGVL, AGVL, KGVL) and QSP 5 1 0.000 0.2827 0.2877 0.2218 0.145 2 1.0000 0.2827 0.2877 0.2218 0.1145 2 2 1 1.0000 0.3566 0.2335 0.1385 0.0700 6 1 1.0000 0.2698 0.2230 0.1867 0.0743 1 1.0000 0.2839 0.2265 0.1994 0.1044 7 1.0000 0.2839 <td>1.0000 0.4196 0.3703 0.2451 0.1357 ng with LCV, LSK and QSP 4 0.81 1.0000 0.2229 0.0530 0.1324 0.0519 3 0.78 1.0000 0.2229 0.0530 0.1124 0.0519 3 0.78 1.0000 0.2229 0.0530 0.1124 0.0519 3 0.71 1.0000 0.3224 0.2420 0.1853 0.0809 1 0.72 1.0000 0.3183 0.3187 0.2178 0.1176 2 0.62 1.0000 0.4615 0.4672 0.3275 0.1847 Case 4: Clustering on 4 0.000 0.4615 0.4672 0.3275 0.1847 Case 4: Clustering on 4 1.0000 0.2375 0.1054 0.1379 0.0527 4 0.72 1.0000 0.23179 0.1054 0.1379 0.0527 4 0.72 1.0000 0.2327 0.2877 0.2218 0.1415 2 0.72 1.0000 0.2692 0.2237 0.2877 0.2268 0.276</td> <td>1.0000 0.4196 0.3703 0.2451 0.1359 ng with LCV, LSK and QSP 0.0530 0.1324 0.0519 1.0000 1.0000 0.2229 0.0530 0.1324 0.0519 1.0784 1.0000 0.2229 0.0530 0.1324 0.0519 1.0784 5.0.73 1.0000 0.3224 0.2420 0.1853 0.0809 1.0.72 0.47 1.0000 0.3183 0.3187 0.2178 0.1176 2.0.62 0.62 1.0000 0.4615 0.4672 0.3275 0.1847 Case 4: Clustering on GEV parameters ng on GEV parameters (HGVL, AGVL, KGVL) and QSP 5 0.72 0.46 1.0000 0.2375 0.1054 0.1379 0.0527 4 0.77 0.39 1.0000 0.2317 0.1054 0.1379 0.0527 4 0.72 0.48 1.0000 0.2317 0.2077 0.2218 0.1145 2 0.72 0.48 1.0000 0.3216 0.2319 0.0952 3 0.70 6 0.76 0.41 <td>1.0000 0.4136 0.3703 0.2451 0.1359 ng with LCV, ISK and QSP 4 0.81 0.32 0.85 1.0000 0.2229 0.0530 0.1324 0.0519 3 0.78 0.38 0.77 1.0000 0.2229 0.0530 0.1324 0.0519 3 0.78 0.38 0.77 1.0000 0.3224 0.2420 0.1853 0.0809 1 0.72 0.49 0.68 1.0000 0.3224 0.4672 0.3275 0.1847 0.62 0.67 0.52 1.0000 0.4615 0.4672 0.3275 0.1847 Case 4: Clustering on GEV parameters (MGVL, AGVL, AGVL, KGVL) and QSP 5 0.80 0.34 0.82 1.0000 0.2375 0.1054 0.1379 0.0527 4 0.77 0.39 0.80 1.0000 0.2375 0.1054 0.1379 0.0527 1 0.76 0.41 0.74 1.0000 0.2370 0.1265 0.1319 0.9952 3 0.70 0.51 0.68 1.0000 0.2690</td><td>1.0000 0.1196 0.3703 0.2451 0.1159 ng with LCV, LSK and QSP 0.0530 0.1124 0.0519 3 0.784 0.310 0.322 0.85 0.371 1.0000 0.2229 0.0530 0.1124 0.0519 3 0.784 0.310 0.71 0.45 1.0000 0.2240 0.2420 0.1853 0.0009 0.622 0.67 0.71 0.42 1.0000 0.3224 0.2420 0.1375 0.1176 2 0.62 0.67 0.52 0.38 1.0000 0.4615 0.4672 0.3275 0.1847 0.622 0.67 0.52 0.38 1.0000 0.2375 0.1054 0.1379 0.0527 1 0.76 0.41 0.74 0.31 1.0000 0.2375 0.1239 0.0644 0.72 0.48 0.46 0.31 1.0000 0.2375 0.1239 0.0700 6 0.72 0.48 0.68 0.47 1.0000 0.2695 0.2216 0.1303 0.3216 0.2319 0.2605</td></td>	1.0000 0.4196 0.3703 0.2451 0.1357 ng with LCV, LSK and QSP 4 0.81 1.0000 0.2229 0.0530 0.1324 0.0519 3 0.78 1.0000 0.2229 0.0530 0.1124 0.0519 3 0.78 1.0000 0.2229 0.0530 0.1124 0.0519 3 0.71 1.0000 0.3224 0.2420 0.1853 0.0809 1 0.72 1.0000 0.3183 0.3187 0.2178 0.1176 2 0.62 1.0000 0.4615 0.4672 0.3275 0.1847 Case 4: Clustering on 4 0.000 0.4615 0.4672 0.3275 0.1847 Case 4: Clustering on 4 1.0000 0.2375 0.1054 0.1379 0.0527 4 0.72 1.0000 0.23179 0.1054 0.1379 0.0527 4 0.72 1.0000 0.2327 0.2877 0.2218 0.1415 2 0.72 1.0000 0.2692 0.2237 0.2877 0.2268 0.276	1.0000 0.4196 0.3703 0.2451 0.1359 ng with LCV, LSK and QSP 0.0530 0.1324 0.0519 1.0000 1.0000 0.2229 0.0530 0.1324 0.0519 1.0784 1.0000 0.2229 0.0530 0.1324 0.0519 1.0784 5.0.73 1.0000 0.3224 0.2420 0.1853 0.0809 1.0.72 0.47 1.0000 0.3183 0.3187 0.2178 0.1176 2.0.62 0.62 1.0000 0.4615 0.4672 0.3275 0.1847 Case 4: Clustering on GEV parameters ng on GEV parameters (HGVL, AGVL, KGVL) and QSP 5 0.72 0.46 1.0000 0.2375 0.1054 0.1379 0.0527 4 0.77 0.39 1.0000 0.2317 0.1054 0.1379 0.0527 4 0.72 0.48 1.0000 0.2317 0.2077 0.2218 0.1145 2 0.72 0.48 1.0000 0.3216 0.2319 0.0952 3 0.70 6 0.76 0.41 <td>1.0000 0.4136 0.3703 0.2451 0.1359 ng with LCV, ISK and QSP 4 0.81 0.32 0.85 1.0000 0.2229 0.0530 0.1324 0.0519 3 0.78 0.38 0.77 1.0000 0.2229 0.0530 0.1324 0.0519 3 0.78 0.38 0.77 1.0000 0.3224 0.2420 0.1853 0.0809 1 0.72 0.49 0.68 1.0000 0.3224 0.4672 0.3275 0.1847 0.62 0.67 0.52 1.0000 0.4615 0.4672 0.3275 0.1847 Case 4: Clustering on GEV parameters (MGVL, AGVL, AGVL, KGVL) and QSP 5 0.80 0.34 0.82 1.0000 0.2375 0.1054 0.1379 0.0527 4 0.77 0.39 0.80 1.0000 0.2375 0.1054 0.1379 0.0527 1 0.76 0.41 0.74 1.0000 0.2370 0.1265 0.1319 0.9952 3 0.70 0.51 0.68 1.0000 0.2690</td> <td>1.0000 0.1196 0.3703 0.2451 0.1159 ng with LCV, LSK and QSP 0.0530 0.1124 0.0519 3 0.784 0.310 0.322 0.85 0.371 1.0000 0.2229 0.0530 0.1124 0.0519 3 0.784 0.310 0.71 0.45 1.0000 0.2240 0.2420 0.1853 0.0009 0.622 0.67 0.71 0.42 1.0000 0.3224 0.2420 0.1375 0.1176 2 0.62 0.67 0.52 0.38 1.0000 0.4615 0.4672 0.3275 0.1847 0.622 0.67 0.52 0.38 1.0000 0.2375 0.1054 0.1379 0.0527 1 0.76 0.41 0.74 0.31 1.0000 0.2375 0.1239 0.0644 0.72 0.48 0.46 0.31 1.0000 0.2375 0.1239 0.0700 6 0.72 0.48 0.68 0.47 1.0000 0.2695 0.2216 0.1303 0.3216 0.2319 0.2605</td>	1.0000 0.4136 0.3703 0.2451 0.1359 ng with LCV, ISK and QSP 4 0.81 0.32 0.85 1.0000 0.2229 0.0530 0.1324 0.0519 3 0.78 0.38 0.77 1.0000 0.2229 0.0530 0.1324 0.0519 3 0.78 0.38 0.77 1.0000 0.3224 0.2420 0.1853 0.0809 1 0.72 0.49 0.68 1.0000 0.3224 0.4672 0.3275 0.1847 0.62 0.67 0.52 1.0000 0.4615 0.4672 0.3275 0.1847 Case 4: Clustering on GEV parameters (MGVL, AGVL, AGVL, KGVL) and QSP 5 0.80 0.34 0.82 1.0000 0.2375 0.1054 0.1379 0.0527 4 0.77 0.39 0.80 1.0000 0.2375 0.1054 0.1379 0.0527 1 0.76 0.41 0.74 1.0000 0.2370 0.1265 0.1319 0.9952 3 0.70 0.51 0.68 1.0000 0.2690	1.0000 0.1196 0.3703 0.2451 0.1159 ng with LCV, LSK and QSP 0.0530 0.1124 0.0519 3 0.784 0.310 0.322 0.85 0.371 1.0000 0.2229 0.0530 0.1124 0.0519 3 0.784 0.310 0.71 0.45 1.0000 0.2240 0.2420 0.1853 0.0009 0.622 0.67 0.71 0.42 1.0000 0.3224 0.2420 0.1375 0.1176 2 0.62 0.67 0.52 0.38 1.0000 0.4615 0.4672 0.3275 0.1847 0.622 0.67 0.52 0.38 1.0000 0.2375 0.1054 0.1379 0.0527 1 0.76 0.41 0.74 0.31 1.0000 0.2375 0.1239 0.0644 0.72 0.48 0.46 0.31 1.0000 0.2375 0.1239 0.0700 6 0.72 0.48 0.68 0.47 1.0000 0.2695 0.2216 0.1303 0.3216 0.2319 0.2605

The shapes of the regionalized frequency growth curves for cluster regions (not the actual cluster region numbers) depends on the clustering variables and the underlying probability distribution used. For example, for the EV1 distribution, the regionalized frequency growth curves are different when clustering on LCV and QSP (Case 1) when compared to clustering on LCV, LSK and QSP (Case 3). The differences are more prominent with the GEV probability distribution. It is clear from Figures 3.14-3.21 that EV1 distribution produces straight linear graphs with normalized discharge ratios ranging from 0.0-5.0 since it has only two parameters (as defined by the coefficient of variation, This distribution would be appropriate for flood data LCV). exhibiting a moderate skew close to the EV1 skew of 1.14. In contrast, the GEV distribution produces pronounced non-linear curves with normalized discharge ratios ranging from 0.0-10.0 since it has an additional parameter to capture high skew commonly present in the flood data (as defined by the coefficient of skewness, LSK). Thus, the three-parameter GEV distribution is able to model the upper tail (return periods greater than 20 years) better than the two-parameter EV1 distribution for highly skewed flood data. This is clearly evident for regions that have steeper regionalized flood frequency growth curves, and, hence are characterized by high coefficients of variation, LCV, and skewness, LSK. For example, in Figures 3.14 and 3.15 for Case 1 (cluster regions delineated using variables LCV and QSP) cluster region number 3 has the flattest curve with regionalized EV1 parameters of MEVL = 0.80 (location) and AEVL = 0.34 (scale) and regionalized GEV parameters MGVL = 0.80 (location), AGVL = 0.34 (scale) and KGVL = -0.01Since the shape parameter, KGVL, of the GEV (shape). distribution is close to zero, the EV1 and GEV flood frequency curves for this cluster region are similar. However, a comparison of the regionalized flood frequency curves for the steepest curves associated with cluster

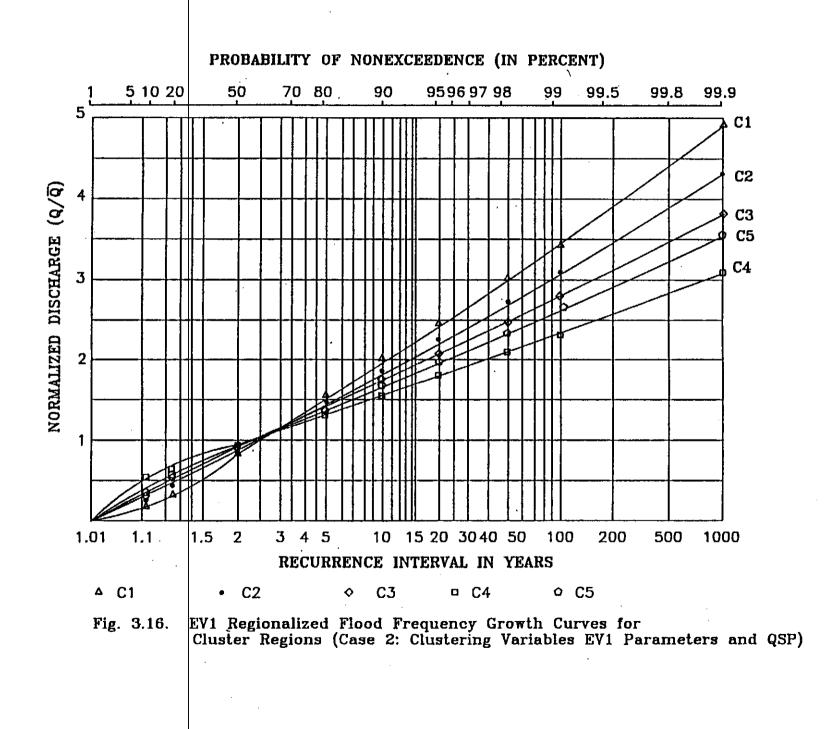
region number 5 (having regionalized parameters for EV1 : MEVL = 0.63; AEVL = 0.64 and for GEV : MGVL = 0.55; AGVL = 0.42 ; KGVL = -0.33) shows considerable difference in the normalized discharge values for return periods greater than 20 years. In all clustering cases, the regionalized flood frequency growth curves are distinct between regions indicating a successful delineation of flood regions (homogeneous within but distinct from other regions) using cluster analysis.

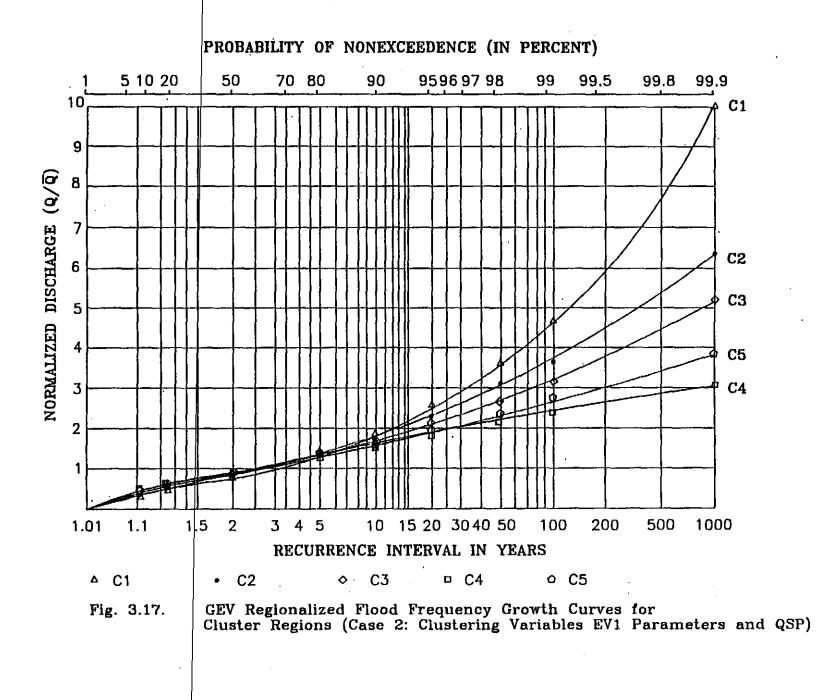
An examination of the regionalized flood frequency curves for the USGS regions, as illustrated in Figures 3.22 and 3.23, shows very little difference between the regions for both EV1 and GEV probability distributions. In both cases, the normalized flood discharge values range from 0.0-5.0 similar to the EV1 distribution for the cluster regions. Thus, at least in terms of their flood frequency growth curves, the USGS regions show more homogeneity across regions than the cluster regions. Furthermore, the frequency growth curves are not very steep for all the seven USGS regions (GEV shape parameter ranges from 0.0 to -0.17 as shown in Table 3.13).

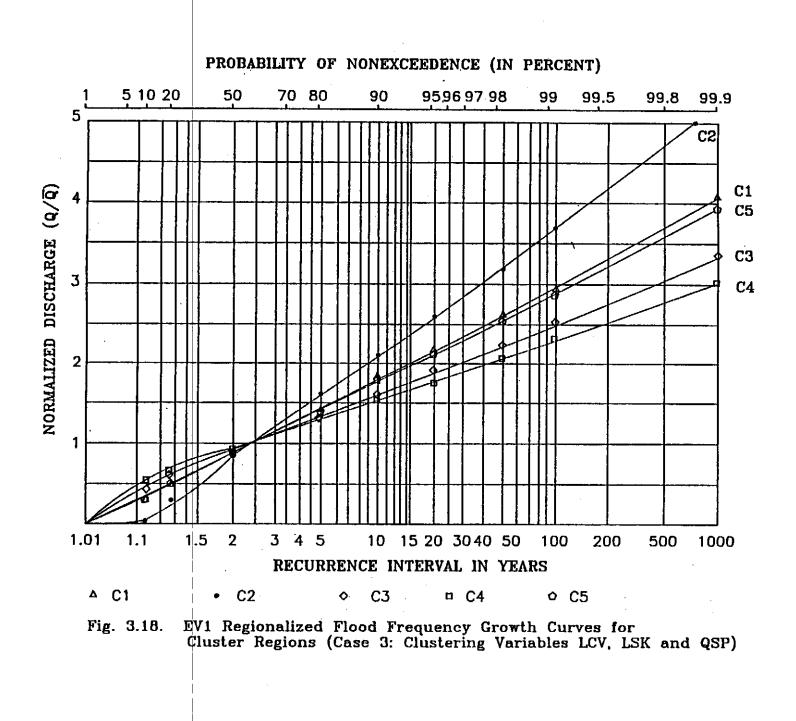


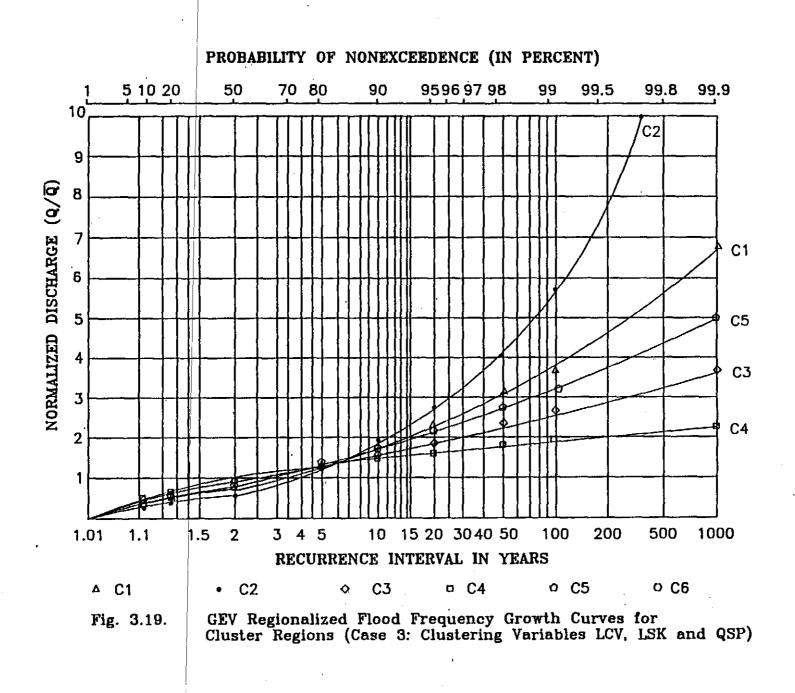


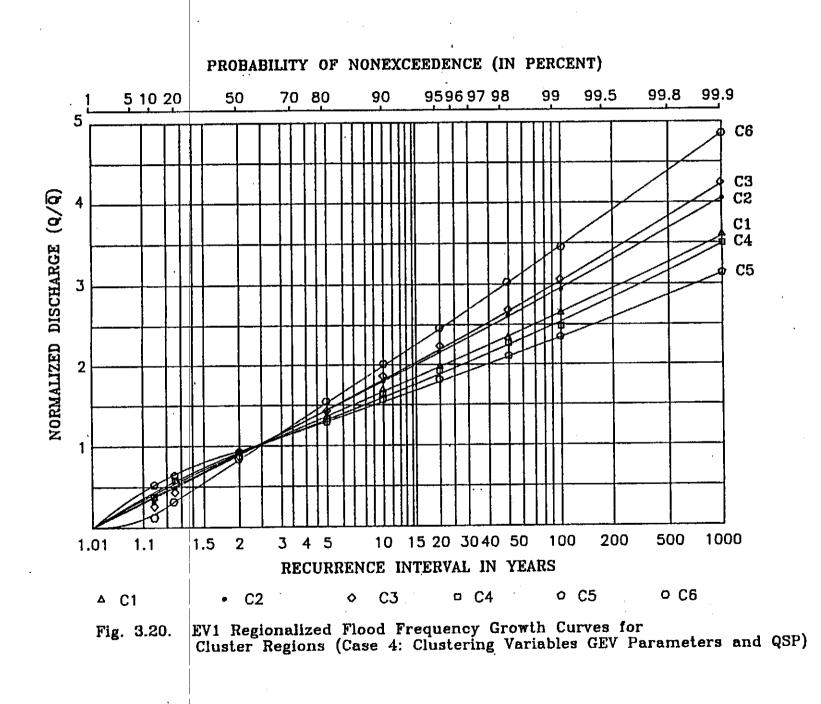
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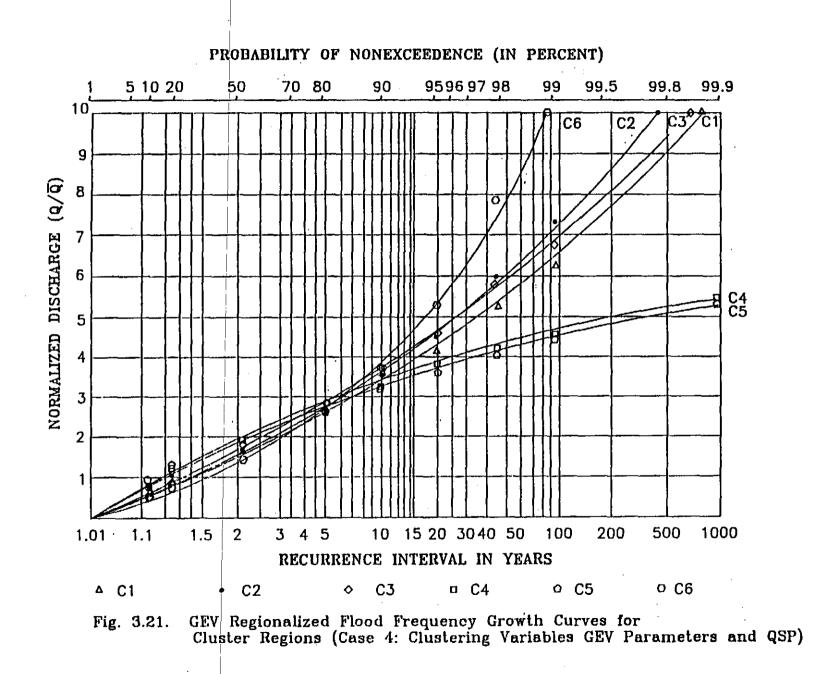


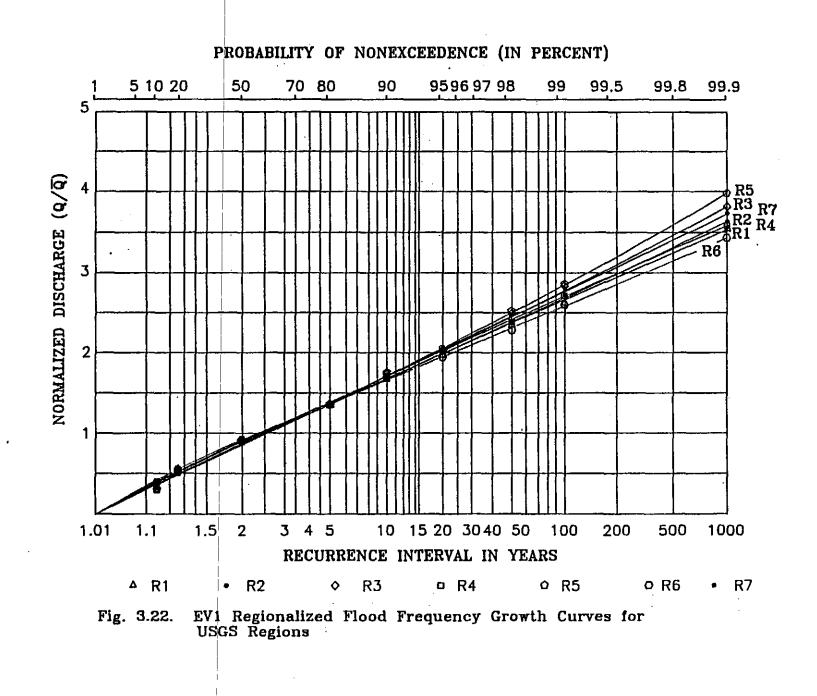


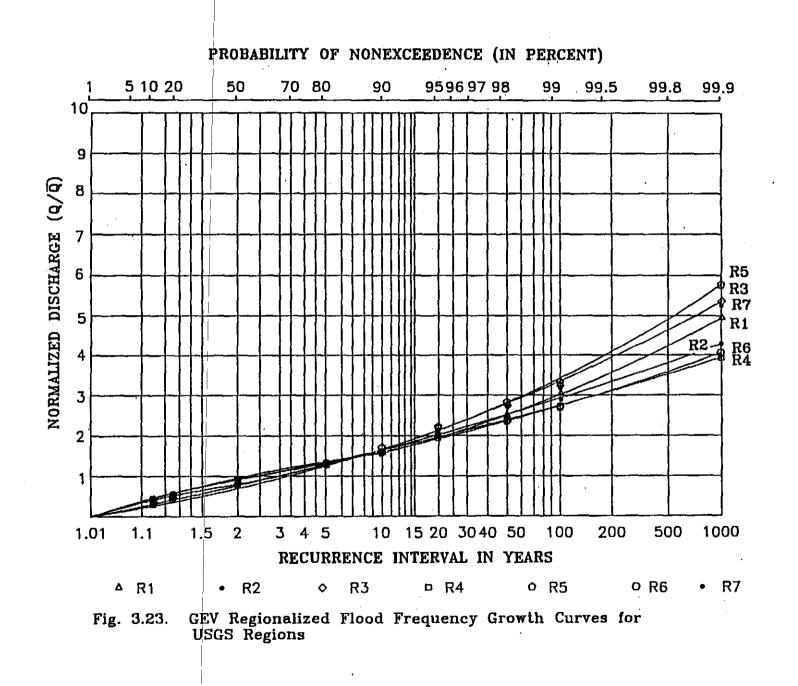












VERFICATION AND COMPARISON OF CLUSTER AND USGS FLOOD REGIONS

a) Hydrologic Characteristics of Flood Regions: The presence of a high degree heterogeneity (or the lack of homogeneity) in the flood characteristics between gauged sites within flood regions, as measured by important statistical properties of the AMF series observed at the site, can adversely affect the benefits derived from regionalization. Ideally, one would like to delineate flood regions that are homogeneous within themselves but distinct from others. As pointed out by Lettenmaier et al (1987), an implicit assumption of most index-flood methods of regionalization, similar to the one used in this study, is that the regions are homogeneous. This would imply that statistical moment ratios of the AMF series, like the coefficient of variation, LCV or CV (both measure the scale of a flood frequency distribution and are closely related), are identical at each of the gaged sites within a region. In reality this will never be the case. With this in view, Lettenmaier et al (1987) examined the effects of heterogeneity of the coefficient of variation on various flood regionalization schemes in conjunction with several parent flood probability distribution. They observed that the advantage of using any regionalization method is reduced for large values of regional average mean coefficient of variation, M(CV), and the range, R(CV), of the values of the coefficient of variation of flood data at each of the gaged sites within the region. Thus, these and similar studies clearly indicate the importance of observing the statistical trends of variables controlling flood response within flood regions.

With the above discussion in mind, statistical trends of important hydrologic characteristics, as measured each of the gauged sites within each of the cluster and USGS regions, are developed and examined in detail. For the four clustering cases (Case 1-Case 4), these trends are

illustrated in Tables 3.14-3.17. Specifically, trends in the mean, median, maximum, minimum and range statistics of clustering variables (L-moments, QSP and parameters), watershed physical characteristics and other hydrologic variables are included in these tables. The cluster regions in each table are arranged in the order of increasing steepness (i.e. increasing coefficients of variation, LCV and/or skewness, LSK) of the regionalized flood frequency growth curve representing each region (refer to previous section).

The trends in the mean and median values of the clustering variables like the L-moments, parameters of the probability distribution and QSP are quite obvious since cluster analysis will group these variables into regions having small, medium to large values. For instance, Table 3.14(a) shows a clear and distinct mean and median values of regional average L-moment ratios (LCV, LSK and LKUR) and the conventional method of moment ratios (CV, SK and KUR) when clustering with LCV and QSP (Case 1). A similar trend is observed for clustering Cases 2-4 as well.

Table 3.19 shows the variation of the regional median values of the coefficient of variation, M(LCV) / M(CV), including its range within each region, R(LCV) / R(CV). The median value of the coefficient of variation, M(LCV), varies from 0.241-0.434 over the five cluster regions for Case 1 and 0.228-0.467 for Case 2. The trend in the median coefficient of variation, M(CV), (as estimated from the method of moments) varies from 0.438-0.936 for Case 1 and 0.375-1.035 for Case 2. A comparison of M(LCV) and M(CV) for other clustering cases shows similar variation. For all cases, M(LCV) and M(CV) are less than 0.467 and 1.035, respectively. The range in the regional median coefficients of variation, R(LCV) and R(CV), vary from 0.087-0.201 and 0.181-0.725, respectively, over all clustering cases. Thus, each cluster region is fairly homogeneous with respect to the variation of the regional median coefficient of

variation. The differences of all regional mean and median L-moments (LCV and LSK in particular), ranging from small to large, make the cluster regions distinct from one another. It is for this reason, as discussed in the previous section, the cluster regions delineated for the four cases in this study are each associated with a distinct regionalized flood frequency growth curve.

Variation in the mean and median values of other physical characteristics, as illustrated in Tables 3.14(b)-3.17(b), suggests that cluster regions for all four cases (Cases 1-4) are grouped into areas having low, medium or high mean annual flood response. Since drainage area, A_{a} , is highly correlated with the mean annual flood, Q, it follows a similar trend. Thus, the flood regions delineated have either predominantly small, medium or large watersheds. It is interesting to see that the clustering variable QSP (the specific mean annual flood) shows a reverse trend since it decreases with increasing watershed size. In other words small watersheds tend to generate a greater magnitude of direct runoff per unit area than do larger watersheds. The trends in main channel length, L_c , and slope, S_c , and watershed or basin length, B_c, and slope, B_s, show similar trends as the watershed drainage area, A_c, since they are directly proportional to it. Finally, the watershed shape index, B_s, and main channel sinuousity, S_s, do not show a significant trend between cluster regions for obvious reasons. These two dimensionless variables are ratios of quantities having similar magnitudes, either small or large.

An examination of the maximum and minimum values (range is the difference) of all the hydrologic variables (refer to the third and fourth rows of Tables 3.14(b)-3.17(b) for each cluster region) shows some overlap between cluster regions. For example, cluster region 3 for Case 1 (refer to Table 3.14(b)) contains generally the larger watersheds (a mean and median of 203.6 and 104.0 square miles, respectively) with a maximum watershed size of 960.0 square miles.

However, a minimum watershed size of 0.2 square miles indicates the presence of some small watersheds as well. Since flood response is not entirely a function of watershed size but depends on other physical and climatic factors, these small watersheds are incorporated in cluster region 3 because of the small coefficient of variation, LCV, associated with the floods produced. The presence of this overlap between cluster regions is one of the reasons why the ability to discriminate between them based on physical attributes is not very high. This is demonstrated later in section (c).

Table 3.18 shows the trends in the hydrological characteristics of USGS regions and is used to compare similar variables between cluster and the the USGS regions. A noticeable difference exists in the variation of M(LCV) and M(CV) between USGS regions and cluster regions (refer to Table 3.19). For example, the regional median coefficients of variation, M(LCV) and M(CV), are quite uniform (varying from 0.248-0.321 and 0.443-0.617, respectively) between the USGS regions. However, these regions have a larger range values of the median coefficient of variation, R(CV), when compared to cluster regions indicating a diversity of watersheds (small to large LCV and CV) contained within each region.

An examination of the mean and median values of the contributing drainage area, A_c , (associated with each of the gauged sites within a region) suggests that the USGS regions have fairly uniform distribution of small to large watersheds within their regions. A similar trend is observed with the mean annual flood, \overline{Q} , since this variable is highly correlated with the contributing drainage area. In contrast, cluster analysis tends to produce regions that have either predominantly small, medium or large watersheds (refer to Tables 3.14(b)-3.17(b)). In this context, it must be emphasized that small watersheds having low mean annual flood, Q, are, generally, associated with high LCV and/or

LSK values while using clustering schemes that included the latter variables as clustering variables.

The distribution of the mean and median values of watershed characteristics such as main channel sinuousity, S_s , and basin shape, both of which involve ratios of similar magnitudes (i.e. either small or large), show similar differences between cluster and USGS regions. Main channel and basin length follow the same trend as the contributing drainage area, A_c , since these variables are highly correlated to it.

An examination of the ranges of the median values of the hydrologic characteristics discussed above indicates that, with the exception of main channel sinuousity, S_s , and basin shape index, B_s (which remain similar for reasons stated in the previous paragraph), the hydrologic characteristics across all cluster regions show more variability than the USGS regions. This is particularly an important asset for discriminating between regions, and, as illustrated later, is the main reason why the USGS regions can not be discriminated easily. TABLE 3.14(a). Comparison of Important Statistics of Regional Moment Ratics Using L-Moments and Conventional Method of Moments: Clustering with LC7 and QSP **

		L	-Noments	*	Xet	hod of Ho	<u>ents</u>
<u>Region</u> No:	<u>No: of</u> <u>Sites</u>	LCY	LSK	LKOR	2	<u>51</u>	KUR
Mean /)	nedian /	<u>max / m</u>	in / ran	<u>ae</u>			
3	93	0.238	0.176	0.181	0.421	0.811	1.628
		0.241	0.154	0.181	0.438	0.721	0.526
		0.282	0.408	0.468	0.709	4.689	26.467
		0.131	-0.130	0.038	0.202	-0.950	-1.636
		0.151	0.538	0.430	0.507	5.639	26.103
-4	30	0.282	0.199	0.184	0.489	0.606	0.520
		0.263	0.148	0.161	0.485	0.428	-0.027
		0.334	0.435	0.317	0.673	2.511	7.464
		0.221	-0.072	-0.021	0.366	-0.725	-2.348
		0.113	0.507	0.338	0.307	3.239	9.812
1	89	0.324	0.276	0.191	0.623	1.390	2.556
		0.318	0.296	0.182	0.618	1.274	1.615
		0.408	0.540	0.384	0.908	3.902	19.933
		0.285	-0.056	-0.008	0.486	-0.455	-1.595
		0.123	0.596	0.392	0.422	4.357	21.528
2	16	0.386	0.306	0.190	0.738	_ 1.182	1.543
		0.386	0.280	0.169	0.701	1.122	0.412
		0.467	0.473	0.500	0.941	3.136	10.954
		0.338	0.102	0.021	0.612	0.131	-1.424
		0.129	0.371	0.479	0.329	3.005	12.378
5	25	0.443	0.404	0.278	0.962	2.077	5.035
		0.434	0.405	0.304	0.936	2.155	5.122
		0.510	0.614	0.475	1.376	3.551	14.323
		0.401	0.190	0.045	0.732	0.471	-0.845
		0.129	0.424	0.430	0.644	3.080	15.168

** Regions arranged in increasing steepness of the corresponding flood frequency growth curves (i.e. increasing LCV or LSK)

* Regional averages of L-moment ratios are weighted by the number of years of record at each site within each region. Conventional. moment ratios are simple arithmetic averages.

f The coefficient of kurtosis, XUR, is computed relative to the normal probability distribution which has a XUR = 3.0. Therefore, observed kurtosis is obtained by adding a value of 3.0.

TABLE 3.14(b). Comparison of Important Statistics of Regional Hydrologic Characteristics: Clustering with LCV and QSP **

Reg.	8	***	QSP	₽_#	B 1	LG	s.	°.,	X
¥o.	(cfs)	(sq. mi.)	(csa))	(11)	(mi)	(4)		(YES)
Sean	/ Media	1 / <u>2017</u> /	<u>nin</u>						
3	8539.6	203.6	\$6.3	2.4	18.7	33.8	0.59	1.6	29.2
•	5858.4	104.0	61.6	2.3	16.8	24.9	0.22	1.5	30
	31384.4	960.0	253.1	5.8	66.0	106.9	8.28	3.1	68
	35.2	0.2	14.6	0.2	0.8	1.1	0.05	1.0	7
-4		1.2 -	430.6	2.3-	1.5	-1.7	1.78	1.2	13.1
		0.8							
	2377.1	5.6	\$92.7	6.1	3.9	4.4	4.66	1.9	34
	48.6	0.1	275.8	0.9	0.5	0.6	0.53	1.0	.7
1	6836.6	152.2	92.4	2.3	15.6	26.2	0.66	1.5	27.1
	4752.0	65.8	73.2	2.2	13.2	19.0	0.15	1.5	27
		936.0	251.0	4.3	56.2	102.5	3.83	2.8	- 63
	40.8	0.2	18.0	Q.7	0.6	0. 7	0.04	1.0	7
2	414.1	1.1	448.5	1.5	1.4	1.8	2.53	1.3	13.4
	220.4	.0.6	384.7	1.4	0.9	1.2	2.14	1.3	10.5
		7.8	823.9	2.4	7.2	9.2	6.49	2.0	30
	107.1	0.1	290.5	0.6	0.4	0.6	0.20	1.1	
5		26.1							17.4
	382.4	1.4	163.7	1.9	2.4	2.9	1.40		15
	16781.4	246.0	368.4	3.6	23.1	58.7	9.66	2.5	32
	67.1	0.6	57.5	0.3	0.5	0.7	0.10	1.0	· 9

** Regions arranged in increasing steepness of the corre-sponding flood frequency growth curves (i.e. increasing LCV or LSKEW).

TABLE 3.15(a). Comparison of Important Statistics of Regional Moment Ratios Using L-Moments and Conventional Method of Moments: Clustering with EV1 Parameters and CSP **

		L	-Howents	•	Met	hod of Ho	ments
Region No:	<u>No: of</u> Sites	LCY	LSK	LXUR	S	SX	KUR
fean / I	Median /	<u>ax / a</u>	in				
4	79	0.226	0.142	0.182	0.405	0.754	1.537
		0.232	0.143	0.181	0.412	0.695	0.422
		0.255	0.408	0.468	0.709	4.689	21.467
		0.131	-0.130	0.038	0.202	-0.950	-1.636
		0.137	0.53#	0.430	0.507	5.639	23.103
5	30	0.276	0.165	0.161	0.485	0.591	0.455
		0.277	0.147	0.161	0.482	0,380	-0.050
		0.334	0.435	0.317	0.673	2.511	7.464
		0.221	-0.072	-0.021	0,366	-0.728	-2.348
		0.113	0.507	0.338	0.307	3.239	9.812
3	91	0.314	0.260	0.182	0.592	1.340	2.542
		0.310	0,262	0.179	0.586	1.265	1.615
		0.374	0.540	0.384	0.852	3.902	19:933
		0.270	-0.056	-0.008	0.446	-0.455	-1.595
		0.104	0.596	0.392	0.406	4.357	21.528
2	10	0_374	0.288	0.206	0.716	1.286	1.963
		0.379	0.250	0.205	0.701	1.133	1.089
		0.426	0.473	0.500	0.928	3.136	10.954
		0.306	0.102	0.021	0.547	0.131	-1.424
		0.120	0.371	0.479	0.381	3.005	12.378
1	43	0.425	6.370	0.237	0.872	1.768	3.873
		0.410	0.366	0.212	0.808	1.662	2.887
		0.530	0.614	0.475	1.176	3.551	14.323
		0.371	0.143	0.033	0.651	0.306	-1.285
		0.159	0.471	0.442	0.725	3.245	15.608

** Regions arranged in increasing steepness of the corresponding flood frequency growth curves (i.e. increasing LCV or LSK).

 Regional averages of L-moment ratios are weighted by the number of years of record at each site within each region. Conventional moment ratios are simple arithmetic averages.

/ The coefficient of kurtosis, KUR, is computed relative to the normal probability distribution which has a KUR = 3.0. Therefore, observed kurtosis is obtained by adding a value of 3.0.

. . .

TABLE 1.15(b). Comparison of Important Statistics of Regional Bydrologic Characteristics: Clustering with EV1 Parameters and QSP ++

Reg.	. Q	л е	QS 2	в,	B 1	Lc	s _c	5 <u>.</u>	X
No.	(cfs)	(sq. mi.)	(CSR)		(mi)	(mi)	(\$)		(yzs)
Mean	/ Media	1 Jax./	ain						
4	\$518.7	211.0	85.3	2.4	18.8	35.0	9.56	1.7	29.0
		104.0							
	31384.4	960.0							
	35.2	0.3	14.6	0.2	0.8	1.1	0.05	1.0	7
5	468.5	1.1	416.8	2.4	1.5	1.7	1.75	1.2	12.5
	262.4	0.7	398.5	2.1	1.4	1.7	1.47	1.1	10
	2377.1	1.0	\$92.6	6.1	1.9	4.4	4.66	1.5	34
	38.6	0.1	241.3	1.1	0.5	0.6	0.53	1.0	7
3		153.2							
		82.3							
		\$36.0							
	40.8	0.2	18.0	0.7	0.6	0.7	0.04	1.0	7
2		. 0.5							
		· 0.4	522.5	1-1	0.7	0.9	2.44	1.3	10.5
	655.3	1.0	823.9	2.4	1.0	1.9	5.47	2.0	26
	107.1	0.1	377.6	9.5	0.4	9.6	0.20	1.1	
` 1		59.1							
	537.1								
		936.0				\$9.6			
	67.1	0.5	27.0	0.3	0.5	0.7	0.08	1.0	

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** Regions arranged in increasing steepness of the corre-sponding flood frequency growth curves (i.e. increasing LCV or LSKEW). TABLE 3.16(a). Comparison of Important Statistics of Regional Moment Ratios Using L-Moments and Conventional Method of Moments: Clustering with LCV,LSK and QSP **

		L	Homents		Xet	hod of Nor	tents
Region Ng:	<u>No: of</u> Sites	1451.	LSK	LXUR	a	SI.	KUR [#]
tean / 1	median /	1 x / x	in / ran	29			
4	44	0.223	0.053	0.132	0.380	0.071	-0.307
-		0.228	0.058	0.135	0.375	0.131	-0.388
		0.300	0.190	0.380	0.516	1.346	3.917
		0.131	-0.130	-0.021	0.202	-0.950	-2.348
		0.169	0.320	0.401	0.718	2.296	6.265
3	75	0.261	0.201	0.171	0.476	1.100	1.853
		0.258	0.204	0.183	0.478	1.045	1.313
		0.350	0.345	0.468	0.631	4.423	22.974
		0.180	0.076	-0.001	0.316	0.095	-1.595
		0.170	0.269	0.469	0.315	4.328	24.569
5	38	0.322	0.242	0.185	0.578	0.836	0.936
-		0.306	0.222	0.165	0.566	0.807	0.197
		0.426	0.473	0.500	0.928	3.136	-10.954
		0.225	-0.056	0.021	0.369	-0.582	-1.571
		0.201	0.529	0.479	0.559	3.718	12.525
1	79	0.338	0.319	0.218	0.669	1.640	3.506
+		0.345	0.323	0.197	0.640	1.485	1.903
		0.431	0.540	0.384	0.968	4.689	26.467
		0.249	0.134	0.045	0.502	0.311	-0.958
		0.182	0.406	0.339	0.466	4.378	27.42
2	17	0.462	0.467	0.328	1.054	2.423	6.724
	*'	0.467	0.472	0.343	1.035	2.531	6.77:
		8.530	0.614	0.475	1.376	3.551	14.32
	••	0.393	0.312	0,136	0.824	1.024	-0.48
		0.137	0.302	0.339	0.552	2.527	14.80

** Regions arranged in increasing stampness of the corresponding flood frequency growth curves (i.e. increasing LCV or LSK).

 Regional averages of L-moment ratios are weighted by the number of years of record at each site within each region. Conventional moment ratios are simple arithmetic averages.

The coefficient of kurtosis. KUR, is computed relative to the normal probability distribution which has a KUR = 3.0. Therefore, observed kurtosis is obtained by adding a value of 3.0.

TABLE 3.16(b). Comparison of Important Statistics of Regional Bydrologic Characteristics: Clustering with LCV, LSKEW and QSP **

Reg.	Q	***	QSP	3	^B 1	Le	s _c	°_	×
Xo.	(cfs)	(sq. mi.)) (CIRL)	l	(mi)	(mi)	(\$)		(YIS)
Hean	/ Nedia	1 / 1HX /	nin						
- 4	4355.1	97.8	147.5	2.3	10.8	19.9	1.00	1.5	19.9
	2198.8	17.6	152.2	2.0	7.1	9.8	0.17	1.3	15
	24353.3	745.0	373.5	5.8	37.4	92.6	8.78	2.7	68
	35.2	0.1	22.0	0.2	0.5	0.6	0.05	1.0	7
-3-	-10748.2	- 265.2 -	-48-6	2.5-	23.2	- 40-9 -	0.38	1.7	3412
	8958.5	235.0	47.8	2.3	22.5	38.6	0.16	1.6	34
	31384.4	960.0	222.8	5.5	66.0	106.9	2.95	3.1	63
	198.5	1.1	14.5	0.6	1.7	1.9	0.05	1.0	
5	427.5	1.0 0.6	460.6	2.1	1.3	1.5	2.03	1.2	12.9
	254.1	0.6	423.9	1.8	1.0	1.3	1.88	1.1	10
	24353.3	5.6	\$92.6	-6.1	3.5	4.4	5.87	2.0	34
	35.2	0.1	275.8	0.6	0.4	0.6	0.20	1.0	7
1		108.7							
	3260.0	40.9	84.9	2.0	8.2	13.2	0.45	1.4	- 24
		936.0							
	40.8	0.2	18.4	0.7	0.6	0.7	0.04	1.0	7
2		21.9							
		1.6							
		246.0							
	117.0	0.6	61.8	0.3	0.5	0.7	0.10	1.0	9

** Regions arranged in increasing stampness of the corre-sponding flood frequency growth curves (i.e. increasing LCV or LSKEW).

TABLE 3.17(a). Competison of Important Statistics of Regional Noment Matios Using L-Moments and Conventional Mathod of Moments: Clustering with GEV Persmaters and QSP as

	Sites	Ŋ	151	I.X.D.B	2	35	KUR
Kean /	/ neiber /	Next and	nax / nin / range				
	99	AE5.0	0.006	0.142	0.404	911.0	0.071
,		112.4	111.0	0.141	0.416	0.354	-0.180
		0.100	0.190	0.400	0.537	1.460	7.104
		0.131	-0.110	-0,008	0.202	-0.950	-2.013
		0.169	0-320	0.40 0	0.335	2.810	9.117
4	15	0.271	0.067	0.119	0.455	0.027	-0.652
,	1	0.273	0.052			100.0-	G
		0.108	0.170		•	0.613	1
		0.221	-0.072	-0.021	0.366		-2.348
		0.087	0.242	0.298	111.0	1.541	102.0
		0.286	0.292	0.222	0.552	1.459	1.44.5
,	2	0.293		0.211	0.561	1.425	2.260
		0.368	0.479	0.468	0.052	4.689	26.467
		0.100	0.190	0.100	0.316	263.0	
		0.100	0.289	0.368	0.536	3.994	36.988
~	17	515.0	0.311	922.0		1.421	2.289
ı)	- 61	0.237		0.612	1.275	1.190
	•	0.426	0.473		0.928	3.136	10.954
		0.233	0.205	9:036	0.412	0.484	-1.571
		0.193	0,268	0.464	0.516	2.645	12.525
-	40	0.363	0.215	721.0	0.657	0.913	0.595
•	;	• •	0.224		0.651		i.
		414	0.324	0.223	0.626	1.656	4.772
		0.275	0.076	-0.001	0.506	\$	-1.595
		0.159	0.252	0.224	0.320	1.761	6.367
	92	0.411	0.453	111-0	0.950	2,306	6.087
,	;	0.432		• •		2.221	5.759
		0.510	0.614	0.475	1.376	1.551	14.323
		166.0	0.346	0.160	0.698	1.233	0.346
		0.193	0.268	0.115	0.678	2.316	14.669
an Regions	1 - 5	158	Increasing	ng steepness thereasing	as of the	Corresponding Laki	ding fload
Bed	Regional aver	To sept.	averages of L-moment	ration	velaht	ed by the number	number of
YAAKS	s of record	ord at e	each site	within ea	5	. Conventional	fonal
	nt ration	i ase i ase	aple art	ration are simple writhmetic averages			
	coarticlant of		KUFCOR16, Aleteriter	NUR, 15 COMPUTED			CO Che Sharafara
NOTERL D	al probability		Tanglasel	SEU UDIUA NOTINGIAJSID			TREELOLS

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TABLE 3.17(b). Comparison of Important Statistics of Regional Nydrelogic Charactaristics: Clustering vith GEV Parametars and QSF **

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щ.	(sts)	(mg. mi.)		_	Ē	Ĵ	Ξ		(Jrs)
	/ Nedlan	/ 38% /							
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•	347.9 251.4 2167.0 40.6	5070	485.5 461.4 492.6 302.9	9 H H 9		4460	2.14		10.8 26 26
-	7585.7 5093.6 51384.4	162.2 104.0 960.0	5119 251-0 251-0	4.40	17.9 66.04	29.0	0.00	4440 1101	30.6
~	499.3 296.8 2377.1	1000	453.9 422.2 823.9 275.8	6 T T O		1110			14.6 34
~	3782.5 1245.0 5026.1	729.0 1.1.1	146.2 114.5 401.5 16.0		6 6 6 0 4 0 8 0	14.8 6.1 102.5 0.6	1.33 9.66 0.05	****	55. 55. 56. 56.
•	5125.2 1391.2 25299.4 117.0	101.1 7.8 9.6 0.6	161.0 174.3 378.8 27.0		10.1	5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5	10.0 0.0 0.0 0.0 0.0		10 4 4

A Regions arranged in increasing steepness of the corre-sponding flood frequency growth curves (i.e. increasing LCV or LSKEW).

TABLE J.18(a). Comparison of Important Statiatics of Regional Noment Nation Using L-Noments and Conventional Nethod of Moments for USGS Regions 4a

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Regic No:	on <u>No, of</u> <u>Sites</u>	M(ICA)	-Moments	R [*] (LCV)	Method M(CV)	i of Mome R(CV)	nts ^e R (CV)
<u>no:</u>	<u>91669</u>		<u>KIIKYI</u>		<u>A(CY)</u>	<u>KICVI</u>	<u>K (CV)</u>
		<u>Case</u> 1	: Cluste	cing with	LCV and	<u>OSP</u>	
з	93	0.241	0.151	0.627	0.438	0.507	1.158
4	30	0.283	0.113	0.399	0.485	0.307	0.633
1	89	0.318	0.123	0.387	0.618	0.422	0.683
2	16	0.386	0.129	0.334	0.701	0.329	0.469
5	25	0.434	0.129	0.297	0.936	0.644	0.688
		Case 2: C	lusterin	y with LC	7. LSK ar	d_OSP	
4	44	0.228	0.169	0.741	0.375	0.314	0.837
3	75	0.258	0.170	0.659	0.478	0.315	0.659
5	38	0.306	0.201	0.657	0.566	0.559	0.988
1	79	0.345	0.182	0.528	0.640	0.466	0.728
2	. 17	0.467	0.137	0.293	1.035	0.552	0.533
	Cas	<u>e 3: Clus</u>	<u>stering w</u>	ith <u>Gumbel</u>	<u>l Paramet</u>	<u>ers and</u>	OSP
						•	
4	79	0.232	0.137	0.591	0.412	0.507	1.231
5	30	0.277	0.113	0.408	0.482	0.307	0.637
3 2	91 10	0.310 0.379	0.104	0.335 0.317	0.586		0.693
1	43	0.379	0.120 0.159	0.388	0.701 0.808	0.381 0.725	0.344
-	45		4.133	0.500	0.000	0.725	0.057
	2	<u>ase 4: C</u>]	lustering	with GEV	Paramete	rs and (<u>DSP</u>
5	68	0.241	0.169	0.701	0.416	0.335	0.805
4	15	0.273	0.087	0.319	0.448	0.181	0.404
1 2	81	0.293	0.188	0.642	0.563	0.536	0.952
2	21	0.318	0.193	0.607	0.612	0.516	0.843
3	40	0.356	0.159	0.447	0.653	0.320	0.490
6	28	0.432	0.193	0.447	0.930	0.678	0.729
			<u>USGS</u>	Regions			
6	31	0.248	0.344	1.387	0.443	1.013	2.287
ĭ	32	0.286	0.353	1.234	0.524	0.986	1.882
4	20	0.278	0.217	0.781	0.493	0.463	0.939
2	68	0.293	0.360	1.229	0.521	0.970	1.862
7	38	0.307		0.850	0.573	0.622	1.086
3	26	0.321	0.271	0.844	0.614	0.754	1.228
5	38	0.301	0.370	1.229	0.617	1.131	1.833
					<u>-</u>		
# MI tì	LCV is the regione region and R	nal media LCV is th	n of LCV ne normal	, RLCV is ized regio	the rangonal LCV	e of LC (range/1	/'s for median).
e Mo re	CV is the region and R CV i	al mediar s the nor	n of CV, i malized i	RCV is the regional (e range d IV (range	of CV's : /median)	for the

TABLE 3.19 Comparison of Regional Mean Coefficients of Variation and Their Ranges Within Cluster and USGS Flood Regions

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b) Performance of Regionalized Flood Frequency Models: The performance of the regionalized flood frequency models is evaluated using the following specific criteria in conjunction with Monte Carlo simulation techniques:

- The accuracy of the regional flood frequency model to predict the flood levels associated with different return periods as measured by the bias.
- The precision (as reflected by the overall fit of the model to the flood data) of the flood frequency model as measured by the root mean square error (RMSE).

For each of the cluster regions delineated under the four clustering schemes (Cases 1-4), AMF data is synthetically generated at each of the gauged sites within the region using procedures discussed in Chapter 2. 100 sequences, each having a record length equal to the historic systematic flood record at the gauged site and drawn from both EV1 and GEV populations, are used in the analysis. The regional average L-moments and the corresponding parameters based on synthetically generated flows and the simulation runs compare well with the historical estimates for the flatter regionalized flood frequency growth curves as shown in Tables A.1-A.11, Appendix A. However, the simulated sequences tend to underestimate the higher order L-moments (like LSK and LKUR) with this difference increasing as the regionalized frequency growth curve gets steeper. As pointed below, this is one of the main reasons why the GEV distribution gives larger biases in the flood quantile estimates than the EV1 distribution. The inability of Monte Carlo simulated flood sequences to capture the larger variability associated with historical estimates of higher order moments, like the coefficient of skew, has been widely reported in literature and is referred to as the condition of separation (Matalas, 1975).

The average regional normalized bias and RMSE for select flood quantiles, as estimated using EV1 and GEV flood frequency models, are summarized for the four clustering cases in Tables 3.20-3.23. Similar results for the USGS regions are shown in Table 3.24. The following conclusions are made for the four clustering cases:

- a) As expected, the bias and RMSE generally increase with the return period, T, and with the steepness of the regionalized flood frequency growth curve. The bias changes from positive to negative as the growth curve becomes steeper and, hence, is not uniform across the cluster regions. This is true for both EV1 and GEV distributions over all return periods of interest (10-100 year). Consequently, flood quantiles are overestimated when the growth curves have small slopes and underestimated as the curves become steeper. In a recent study, Landwehr (1980) observed that if the population skew is different (larger or smaller) than the EV1 skew of 1.14, then an EV1 distribution would on the average underestimate the flood quantiles. In this study it appears to hold for a majority of flood regions (particularly those with steep frequency growth curves) indicating regionalized coefficient of skew other than the EV1 skew of 1.14 (refer to Tables 3.14(a)-3.17(a)).
- b) The biases and RMSE for the EV1 flood frequency model are lower than the GEV model for all flood frequency growth curves. However, one would expect the GEV model to do better than the EV1 model, at least in terms of the bias, since it has an additional shape parameter to to better characterize the growth curves, in particular the steep ones. Such is not the case in this study.

The larger biases associated with the GEV are partly due to the condition of separation that exists when using Monte Carlo simulated flood data. In other

words, the use of a three parameter distribution like the GEV may give larger biases than a more parsimonious distribution like the EV1 due the lower variability of the coefficient of skew and higher order moments observed in simulated flood sequences. Furthermore, as pointed by Wallis (1985), the GEV distribution while having a theoretical appeal for fitting flood data, the asymptotic properties on which it is founded may not be satisfied by the small number of independent flood events commonly encountered in practice.

- c) An examination of cluster regions for all four clustering cases (Cases 1-4) indicates that the regional average bias associated with flood levels less than 100 years, ranges from -2.2% to 0.1% for the EV1 distribution and from -14.2% to 0.1% for the GEV distribution while the corresponding RMSE ranges from 9.2% to 21.8% and 9.2% to 43.9%, respectively. These levels are comparable to values reported in previous studies (for example refer to Lettenmaier et al, 1987).
- d) Clustering on the parameters of the probability distribution, as opposed to the L-moments used to estimate them, reduces the bias and RMSE, nominally. This occurs inspite of the fact the shape of the growth curves is affected by the clustering variables used (refer to section on development of flood frequency growth curves).

For the USGS regions the biases and RMSE of the regionalized EV1 and GEV distribution are lower than the cluster regions. This is partly due to the relatively flat regionalized flood frequency growth curves associated with all the seven USGS regions. The regional average bias for all seven regions ranges from -0.9% to 0.1% for the EV1 distribution and -6.0% to 0.0% for the GEV distribution. These biases are usually negative at higher return periods (for example the 100 year) indicating an underestimation of

flood levels. The RMSE ranges from 11.8% to 15.4% for the EV1 distribution and 12.2% to 29.8% for the GEV distribution. Also note that the bias and RMSE are fairly uniform across the seven USGS regions.

A regionalized log-Pearson Type-III distribution (based on L-moments) is not tested in this study since previous studies have clearly shown that EV1 and GEV outperform the log-Pearson Type-III distribution (Wallis, 1985) in estimating flood quantiles. However, since current practice continues to use this distribution, Tables 3.25-3.29 compare log-Pearson Type-III flood quantile estimates (from the USGS method of residuals study using WRC Bulletin 17-B) to the estimates of the EV1 and GEV distributions at select sites within cluster and USGS regions. These sites are chosen to represent gauged sites that have small to large watershed areas, low to high coefficient of variation and skewness associated with the flood data, and the number of years of systematic historic flood records range from 9 to 58 years. Since the true population flood quantile (for a given return period) is unknown, these tables merely serve the purpose of identifying whether flood quantiles are under or over estimated by the recommended regionalized flood frequency distributions in this study. An examination of these tables suggests that, with the exception of using the EV1 distribution at a few sites, flood quantiles are, generally, overestimated when using log-Pearson Type-III distribution.

<u>No:</u>	<u>10 yr.</u>	<u>20 yr</u>	_50 yr.	Probability 100 yr.	<u>Region</u> 1000 vr.
		Σ	lias		
-		_			
					0.006
					-0.002
1					-0.009
2					-0.028
5	0.016	0.002	-0.010	-0.017	-0.030
3	0.004	0.008	0.015	0.020	0.040
4	0.002	0.008			0.061
1	-0.003				-0.016
ž					-0.040
5	0.007	-0.020	-0.053	-0.078	-0.155
		F	MSE		
3	0.094	0.095	0.095	0.095	0.095
4	0.156				0.157
1					0.136
2					0.198
5	0.200	0.197	0.195	0.195	0.194
3	0.095	0.096	0.098	0.100	0.113
					0.227
					0.171
ŝ					
					0.390 0.452
	3 4 1 2 5 3 4 1 2 5 3 4 1 2 5 3	$\begin{array}{cccccccccccccccccccccccccccccccccccc$	E 3 0.003 0.004 4 -0.001 -0.002 1 -0.001 -0.003 2 -0.002 -0.010 5 0.016 0.002 3 0.004 0.008 4 0.002 0.008 1 -0.003 -0.006 2 -0.009 -0.019 5 0.007 -0.020 E 3 0.094 0.095 4 0.156 0.156 1 0.136 0.136 2 0.200 0.199 5 0.200 0.197 3 0.095 0.096 4 0.168 0.172	Bias 3 0.003 0.004 0.005 4 -0.001 -0.002 -0.002 1 -0.001 -0.003 -0.002 1 -0.001 -0.010 -0.017 5 0.016 0.002 -0.010 3 0.004 0.008 0.015 4 0.002 0.008 0.018 1 -0.003 -0.006 -0.009 2 -0.009 -0.019 -0.028 5 0.007 -0.020 -0.053 RMSE 3 0.094 0.095 0.095 4 0.156 0.157 1 0.136 0.136 0.136 2 0.200 0.199 0.199 5 0.200 0.197 0.195 3 0.095 0.096 0.098 4 0.168 0.172 0.179	Bias 3 0.003 0.004 0.005 0.005 4 -0.001 -0.002 -0.002 -0.002 1 -0.001 -0.003 -0.005 -0.007 2 -0.002 -0.010 -0.017 -0.021 5 0.016 0.002 -0.010 -0.017 3 0.004 0.008 0.015 0.020 4 0.002 0.008 0.015 0.020 4 0.002 0.008 0.015 0.020 4 0.002 0.008 0.015 0.020 4 0.002 0.008 0.015 0.020 4 0.003 -0.009 -0.011 0.026 2 -0.009 -0.199 -0.028 -0.033 5 0.007 -0.020 -0.053 -0.078 EMSE 3 0.094 0.095 0.095 0.095 3 0.094 0.095 0.095 0.197 0.198<

TABLE 3.20. Regional Average Normalized Bias and Root Mean Square Error (RMSE) of Quantiles: Clustering with LCV and QSP *

 Regions arranged in increasing stampness of the corresponding flood frequency growth curves (i.e. increasing LCV or LSK)

Probability	Region			<u>Ouantiles</u>	l	-
Distribution	Ng:	<u>10 yr.</u>	<u>20 yr.</u>	<u>50 yr.</u>	<u>100 yr.</u>	<u>1000 yr</u> .
			2	lias		
EV1	4	0.002	0.003	0.004	0.004	0.005
	5	0.002	0.002	0.002	0.002	0.002
	3	0.000	-0.002	-0.004	-0.005	-0.007
	2	0.017	0.011	0.006	0.003	-0.003
	1	0.005	-0.007	-0.017	-0.022	-0.034
GEV	4	0.004	0.009	0.015	0.020	0.039
	5	-0.002	0.004	0.015	0.024	0.061
	3	-0.006	-0.009	-0.013	-0.015	-0.021
	2	-0.015	-0.027	-0.039	-0.047	-0.064
•	1	-0.009	-0.028	-0.053	-0.071	-0.127
			B	MST		
EV1	4	0.092	0.092	0.092	0.092	0.093
	5	0.158	0.158	0.158	0.159	0.159
	3	0.128	0.127	0.127	0.127	0.127
•	2	0.201	0.200	0.199	0.199	0.199
	1	0.193	0.191	0.190	0.189	0.189
SEV	` 4	0.092	0.094	0.096	0.099	0.111
	5	0.163	0.167	0.174	0.183	0.227
	3	0.146	0.146	0.147	0.148	0.155
	2	0.247	0.253	0.265	0.278	0.338
	ī	0.385	0.382	0.380	0.381	0.397

TABLE 3.21. Regional Average Normalized Bias and Root Mean Square Error (RMSE) of Quantiles: Clustering with EVI Parameters and QSP *

 Regions arranged in increasing steepness of the corresponding flood frequency growth curves (i.e. increasing LCV or LSK)

Probability	Decise	·	Qu	antiles		
Distribution	Region No:	<u>10 yr.</u>	<u>20 yr.</u>	<u>50 yr.</u>	100 Vr.	1000 YT.
			8	ias		
EV1	4	0.001	0.003	0.004	0.005	0.006
		0.002	0.002	0.002	0.002	0.002
	3 5 1 2	0.001	-0.002	-0.005	-0.007	-0.010
	1	0.001	-0.003	-0.006	-0.008	-0.012
	2	0.016	0.000	-0.015	-0.022	-0.038
GEV	4	0.014	0.026	0.043	0.056	0.097
	3	0.001	0.003	0.005	0.007	0.015
	1	-0.001	-0.008	-0.018	~0.025	-0.048
	1 5 2	-0.002	-0.003	-0.002	0.000	0.011
	2.	-0.015	-0.054	-0.105	-0.142	-0.259
				MSE		
			-			
EV1	4	0.111	0.111	0.111	0.112	0.112
	3 5 1	0.096	0.096	0.096	0.096	0.096
	5	0.172	0.171	0.171	0.171	0.171
	1	0.146	0.145	0.145	0.145	0.144
	2	0.218	0.214	0.212	0.211	0.210
GEV	4	0.105	0.109	0.116	0.124	0.154
	3	0.098	0.099	0.100	0.101	0.107
	3 1 5	0.188	0.187	0.187	0.188	0.196
	5	0.196	0.198	0.203	0.208	0.233
	2	0.382	0.375	0.375	0.381	0-428

TABLE 1.22. Regional Average Normalized Bias and Root Mean Square Error (RMSE) of Quantiles: Clustering with LCV, LSKEW and QSP \star

* Regions arranged in increasing steepness of the corresponding flood frequency growth curves (i.e. increasing LCV or LSK)

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TABLE 3.23. Regional Average Normalized Bias and Root Mean Square Error (RMSE) of Quantiles: Clustering with GEV Parameters and QSP *

Probability	Region		2 _	wantiles		
Distribution	No:	<u> 10 yr.</u>	_20 YE.	<u>50 yr.</u>	<u>100 yr.</u>	1000 Yr.
			1	liag		
EVI	5	0.003	0.003	0.004	0.004	0.005
	4	0.003	0.003	0.003	0.003	0.003
	1	0.002	0.001	0.000	0.000	-0.001
	ž	-0.009	-0.012	-0.016	-0.018	-0.022
	3	0.005	-0.001	-0.005	-0.008	-0.013
		0.013	0.000	-0.011	-0.017	-0.029
GZV	5	0.009	0.017	0.029	0.038	0.070
	- 4	0,009	0.026	0.051	0.071	0.141
	1	-0.003	-0.007	-0.012	-0.016	0.031
	3	-0.004	-0.004	-0.001	0.003	0.020
	2	-0.008	-0.019	-0.033	-0.044	-0.077
	6	0,016	-0.015	-0.057	-0.089	-0.193
			3	us:		
EV1	5	0.102	0.102	0.103	0.103	0.103
•	4	0.165	0.165	0,166	0.167	0.168
	1	0.114	0.114	0.114	0.114	0.114
	2	0.176	0,176	0.176	0.176	0.176
	3	0.159	0.158	0.157	0.157	0.157
	6	0.202	0.200	0.198	0.197	0.196
GEV	5	9.096	0.098	0.103	0.107	0.127
	4	0.154	0.160	0.173	0.187	0.252
	1	0.130	0.130	0.131	0.132	0.139
	3	0.164	0.165	0.168	0.171	0.189
	2	0.210	0.213	0.219	0.227	0.265
	6	0.439	0.434	0.431	0.433	0.459

Regions arranged in increasing steepness of the corresponding flood frequency growth curves (i.e. increasing LCV or LSK)

TABLE 3.24.	Regional Average Normalized Bias and Root Mean Square
Error (RMSE)	of Quantiles: USGS Regions *

			_ <u> </u>	<u>wantiles</u>		
<u>Probability</u> Distribution	Region No:	<u>10 yr.</u>	<u>20 yr.</u>	<u>50 yr.</u>	<u>100 yr.</u>	1000 vr.
			Bia	E		
EV1	6	-0.001	-0.004	-0.006	-0.008	-0.010
	1	0.014	0.011	0.007	0.006	0.002
	2	0.003	0.001	-0.002	-0.003	-0.006
	4	-0.003	-0.004	-0.006	-0.007	-0.009
	73	0.000	-0.002	-0.004	-0.005	-0.007
	3	0.007	0.004	0.002	0.000	-0.003
	5	0.005	0.001	-0.003	~0.005	-0.009
GEV	6	0.000	-0.002	-0.005	-0.007	-0.010
	7	-0.004	-0.007	-0.010	-0.011	-0.015
	4	0.011	0.017	0.026	0.034	0.066
	2	-0.002	-0.004	-0.006	-0.007	-0.010
	2 1 3	-0.007	-0.014	-0.023	-0.031	~0.055
		0.000	-0.008	-0.017	-0.023	-0.043
	5	-0.010	-0.019	-0.029	-0.037	-0.060
			R	ISE		
EV1	6	0.139	0.138	0.138	0.138	0.138
	ī	0.156	0.155	0.155	0.154	0.154
		0.138	0.137	0.137	0.137	0.137
	2 4	0.118	0.118	0.118	0.118	0.118
	7	0.142	0.142	0.142	0.142	0.142
	3	0.154	0.154	0.154	0.154	0.153
	5	0.152	0.151	0.151	0.150	0.150
GEV	6	0.162	0.163	0.165	0.168	0.185
	7	0.175	0.176	0.179	0.182	0.198
	4	0.122	0.126	0.134	0.142	0.181
	2	0.154	0.154	0.155	0.157	0.164
	ī	0.182	0.182	0.184	0.186	0.202
	3	0.278	0.279	0.281	0.284	0.298
	3 5	0.178	0.179	0.183	0.187	0.209

 Regions arranged in increasing steepness of the corresponding flood frequency growth curves (i.e. increasing LCV or LSK)

TARE 3.25. Comparison of Quantiles at a Yew Selected Stations: Constraint with LTY and OSP (Case 1) 4

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Riartica (avait) Reart of the state of the	Ristion Armit Jac Start Law Core J Law Core J Law Core J Law Law <thlaw< th=""> Law <thlaw< th=""> Law<!--</th--><th>Rinklon Arms Resold Arms Arms</th><th></th><th></th><th></th><th></th><th>PLACE LON</th><th></th><th>9 </th><th></th><th></th><th>aire</th></thlaw<></thlaw<>	Rinklon Arms Resold Arms					PLACE LON		9 			aire
	0 14.1 11 0.11 -0.11 4735 91 5110 0 309-0 50 0.21 0.11 -0.11 4735 91 5110 0 309-0 50 0.21 0.11 -0.11 4735 91 5110 0 309-0 50 0.21 0.11 -0.11 4105 11 1105 0 1 1 0.10 0.13 -0.17 11000 -11 1100 0 0.23 0.10 0.13 -0.17 11000 -11 110 0 0.10 0.13 0.13 0.14 2331 -1 110 0 0.13 0.14 2331 -1 110 -1 110 1 0.13 0.14 2331 -1 110 110 110 1 1 0.14 2331 -1 100 -1 110 1 1 0.14 2331 -1 110 10 10 1 1 0.14 234 -1 100 10 10 1 1 0.14 234 -1 100 10 10 <t< th=""><th>0 14.13 13 0.113 -0.113 1775 57 5310 0 309.0 50 0.23 0.27 34914 -11 38950 0 309.0 50 0.23 0.27 34914 -10 38950 0 309.0 0.23 0.27 34914 -10 38950 - 0 0.59 0.29 0.27 34914 -10 38950 - 0 0.59 0.29 0.77 34914 -10 3130 - 4400 - 4400 - 4400 - 4400 - 4400 - 4400 - 4400 - 4400 - 4400 - 4400 - 4400 - 4400 - 4400 - 440 - 440 - 440 - 440 - 440 - 440 - 440 - 440 - 440 - 440 - 440 - 440 - 440 - 440 - <td< th=""><th>Cluster Region</th><th>station No.</th><th></th><th>Record</th><th>TEL</th><th>155</th><th>(ofe)</th><th>E</th><th>1</th><th>εl</th></td<></th></t<>	0 14.13 13 0.113 -0.113 1775 57 5310 0 309.0 50 0.23 0.27 34914 -11 38950 0 309.0 50 0.23 0.27 34914 -10 38950 0 309.0 0.23 0.27 34914 -10 38950 - 0 0.59 0.29 0.27 34914 -10 38950 - 0 0.59 0.29 0.77 34914 -10 3130 - 4400 - 4400 - 4400 - 4400 - 4400 - 4400 - 4400 - 4400 - 4400 - 4400 - 4400 - 4400 - 4400 - 440 - 440 - 440 - 440 - 440 - 440 - 440 - 440 - 440 - 440 - 440 - 440 - 440 - 440 - <td< th=""><th>Cluster Region</th><th>station No.</th><th></th><th>Record</th><th>TEL</th><th>155</th><th>(ofe)</th><th>E</th><th>1</th><th>εl</th></td<>	Cluster Region	station No.		Record	TEL	155	(ofe)	E	1	εl
	713140 14.3 13 0.13 -0.13 -0.13 17.3 19 5110 131800 309.0 309.0 50 0.231 0.237 34744 -11 3410 131800 309.0 1.7 10 0.231 0.237 34744 -11 3410 130004 1.7 10 0.233 0.237 3400 -1 3410 130004 1.7 10 0.233 -0.077 1100 -1 3410 130004 1.7 10 0.233 -0.077 1100 -1 3410 130014 0.1 0.139 0.139 0.139 -0.13 1100 -1 1113 1301015 0.1 0.139 0.139 0.139 111 111 111 1001015 0.1 0.139 0.141 111	732346 14.3 13 0.13 -0.13 -0.13 173	1104	EXILES.								
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	413500 309.0 50 0.23 0.23 0.23 0.23 0.23 0.24 1111 0.24 1111 0.24 1111 0.24 1111 0.24 1111 0.24 1111	4.13500 309.0 50 0.23 0.21 0.23 0.21 0.23 0.21 0.20 111 1111	•				22.0	5: -	4001	5		
	(13800 309.0 50 0.21 0.27 34746 -10 1111 1111	(13500 309.0 50 0.21 0.27 34746 -11 3850 -21 34746 -11 3850 -21 3460 -21 2104 440 -21 440 -21 440 -21 440 -21 440 -21 440 -21 440 -21 440 -21 440 -21 440 -21 440 -21 140 210					00					
	300041 1.7 10 2.3 2.77 30004 -1.0 4000 31146 0.1 1.0 0.25 -0.07 1110 -1.1 1111 31146 0.1 1.0 0.10 0.43 -0.07 1110 -1.1 1111 31146 0.1 1.9 0.10 0.43 -0.07 1110 -1.1 1111 31146 0.1 1.9 0.10 0.43 -0.07 1110 -1.1 1111 303081 0.1 1.9 0.10 0.43 -0.06 114 -1.1 1111 303081 0.1 0.13 0.10 0.43 -0.06 114 11 1111 303081 0.1 0.13 0.10 0.14 111 111 111 303181 0.4 10 0.10 0.14 111 111 111 303180 0.4 10 0.11 0.11 111 111 111 303180 0.4 10 0.11 0.11 111 111 303180 0.4 10 0.11 0.11 111 111 303180 0.4 10 0.11 0.11	300045 1.7 10 0.25 0.77 39004 -10 2890 300045 1.7 10 0.25 -0.07 1313 -21 14900 301045 0.1 0.25 -0.07 1313 -21 14900 311465 0.1 0.25 -0.07 1313 -21 14900 311465 0.1 0.25 -0.06 141 -1 1313 -1 1313 311465 0.1 0.25 -0.06 141 -1 2134 213 21		435500	309.0	50	0.28	0.27	34746	誯	10980	Ŧ
	300045 1.7 10 0.23 -0.73 1100 -1 1111 31116 0.3 0.55 -0.73 1100 -1 1111 31116 0.3 0.43 233 -1 1113 -1 31116 0.3 0.43 243 233 -1 113 31116 0.3 0.33 0.43 233 -1 113 303043 0.3 0.39 0.43 233 -1 113 303043 0.3 0.33 -0.43 233 -1 113 303043 0.3 0.33 -0.43 233 -1 113 303043 0.3 0.33 -0.46 11 10 10 303043 0.4 0.13 0.41 10 10 10 303043 0.4 0.13 0.14 10 10 10 310345 0.4 0.13 0.11 10 10 10 310345 0.4 0.13 0.13 10 10 10 310345 0.4 0.10 0.13 10 10 10 310345 0.4 0.10 0.13 10 1	300066 1.7 10 0.25 -0.07 1313 -21 1313 31146 0.3 0.43 231 1314 -21 1313 31146 0.3 0.43 231 1314 -21 1313 31146 0.3 0.43 231 -1 1133 31146 0.3 0.43 233 -2 1133 31146 0.3 0.43 233 -1 1133 31146 0.3 0.43 233 -1 114 310395 0.4 11 0.39 -1.43 10 10 310395 0.4 11 0.39 -1.42 10 10 310395 0.4 11 0.39 0.10 114 -1 310395 0.4 11 0.39 0.10 114 10 310395 0.4 11 0.39 -1 10 114 310395 0.4 11 0.39 -1 10 10 310395 0.4 11 0.39 131 10 310395 0.4 10 13 10 10 310396 0.4 0.4 0.4					0.28	0 2	*****		14100	;
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	11146 0.30 -0.703 11106 -71 1110 11146 0.3 0.43 134 -1 111 101011 0.3 0.43 134 -1 111 101011 0.3 0.43 134 -1 111 101011 0.3 0.43 143 -1 111 101013 0.4 10 0.44 141 -1 112 101013 0.4 10 0.44 141 -1 113 101033 0.4 10 0.44 141 -1 114 110133 0.4 11 0.10 0.44 -1 114 110133 0.4 11 0.10 0.41 -1 114 120 110133 0.4 10 0.41 0.10 144 -1 114 110134 0.4 0.4 0.10 144 -1 114 104 110134 0.4 0.4 0.10 144 -1 114 110134 0.4 0.4 0.10 144 -1 114 11014 1.4 0.4 0.4 0.1 14 14 11014 <td>9.131 0.130 -0.73 1100 -1.1330 9.13144 03 043 234 -1.133 234 9.13144 03 043 234 -1.133 234 9.13144 03 043 234 -1.133 234 9.13144 03 043 234 -1.134 234 9.1314 03 043 234 -1.14 234 9.10145 04 11 034 934 -1.14 234 9.10145 04 11 014 234 -900 -1.14 234 9.10145 04 014 2.14 903 -1.141 703 -1.141 9.10146 014 017 9.11 703 -1.141 703 -1.141 703 -1.141 703 -1.141 703 -1.141 703 -1.141 703 -1.141 703 -1.141 703 -1.141 703 -1.141 703 -1.141 703 -1.141 703 -1.141 703 -1.141</td> <td></td> <td>300065</td> <td>1.7</td> <td>97</td> <td>0.23</td> <td>-0.07</td> <td>1312</td> <td>ş</td> <td>1471</td> <td></td>	9.131 0.130 -0.73 1100 -1.1330 9.13144 03 043 234 -1.133 234 9.13144 03 043 234 -1.133 234 9.13144 03 043 234 -1.133 234 9.13144 03 043 234 -1.134 234 9.1314 03 043 234 -1.14 234 9.10145 04 11 034 934 -1.14 234 9.10145 04 11 014 234 -900 -1.14 234 9.10145 04 014 2.14 903 -1.141 703 -1.141 9.10146 014 017 9.11 703 -1.141 703 -1.141 703 -1.141 703 -1.141 703 -1.141 703 -1.141 703 -1.141 703 -1.141 703 -1.141 703 -1.141 703 -1.141 703 -1.141 703 -1.141 703 -1.141		300065	1.7	97	0.23	-0.07	1312	ş	1471	
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TABLE 3.36. Comparison of Quantiles of a Pev Selected Stations: Clustering with EV1 Parameters and OSP (Case 2) ℓ

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1 377076 1.5 10 0.39 0.13 1413 95 1445 254400 13.6 0.49 0.50 872 137 323 254400 13.6 0.49 0.59 1972 137 323 254400 13.6 0.49 0.57 13012 17 3497 254400 13.6 0.49 0.57 14637 2 1410 25400 13.6 0.49 0.51 14637 2 1440 7 1.18 1.11 1.11 1.14 1.4400 - 18400 7 1.18 1.11 1.4400 - 18400 - 18400 7 1.18 1.11 1.1400 - 18400 - 18400 - 18400 - 18400 - 18400 - 18400 - 18400 - 18400 - 18400 - 18400 - 18400 - <t< td=""><td></td><td></td><td></td><td></td><td>0.86</td><td>1.55</td><td></td><td>1</td><td>Ş</td><td>1</td></t<>					0.86	1.55		1	Ş	1
0.39 0.15 1372 1372 1373 254400 13.6 0.50 831 7 347 254400 13.6 0.5 0.5 14637 2 1412 254400 13.6 0.5 0.5 14637 2 1412 254400 13.6 0.5 0.5 14637 2 1410 25450 0.45 0.5 1.1 14600 18400 2645 0.45 0.5 1.1 14600 18400 2745 0.45 0.45 0.5 14600 18400 2745 0.45 0.45 0.5 14600 18400 2745 0.5 1.1 1.1 1.1 1.1 14600 18400 2745 1.1 1.1 1.1 1.1 14600 18400 2745 1.1 1.1 1.1 14600 18400 2745 1.1 1.1 14600 18400 2745 1.1 1.1 14600 18400 2745 1.1 1.1 14600 18400 <t< td=""><td>-</td><td>277070</td><td>1.5</td><td>10</td><td>62.0</td><td>0.15</td><td>1618</td><td>56</td><td>1845</td><td></td></t<>	-	277070	1.5	10	62.0	0.15	1618	56	1845	
234400 11.6 13 0.68 0.50 831 987 234400 11.6 13 0.48 0.57 12013 -17 1369 0.48 0.57 12013 -17 1369 0.48 0.57 14607 - 18400 18400 and a high attion akev, respectively. Redons are arranged in Increasing stammans of the corresponding Redons are arranged in and akenness computed using normalized rev maximum ArS for 100-parson Type-III.					0.39	0.15	1972	5	1252	155
254400 11.6 15 0.45 0.57 12012 -17 1369 The first and second station of anch region were selected based on a and a high attrion akey, respectively. Regions are arranged in increasing steapers of the corresponding flood frequency growth curves [1.6. Increasing LCV of 15K]. Coefficients of variation and services computed using normalized Fay maximum ArS for log-parson Type-III.					0.68	0.50	109	1	547	ł
0.45 0.57 14637 2 10112 The first and second station of each region were selected based on a and a high station atew, respectively. Regions were selected based on a flood frequency growth curves (1.4. increasing icr of the corresponding flood frequency growth curves (1.4. increasing icr of 15K). Coefficients of variation and services computed using increasing for based and a service of some of another using increasing reave maximum ArS for log-parson Type-Til.		254400	3.64	51	0.45	0.57	12012	-17	19961	\$2-
1.18 1.31 1.4400 18400 The first and second station of anch region wars selected based on a and a high station skew, respectively. Regions are arranged in increasing stampness of the corresponding flood frequency growth curves (1.4. increasing LCV or LSR). Conficients of variation and servinas computed using increatized raw maximum AFS for log-paseon Type-LII.				1	0.45	0.57	14637	~	18712	-
The first and second station of arch region were selected based on a and a high station skev, respectively. Regions are stranged in increasing strapases of the corresponding flood frequency growth curves (1.4. increasing ICT of 15K1. Conficients of variation and serves computed using normalized Far maximum AFS for log-parson Type-III.					1.10	1.11	14400	1	18400	ł
Regions are arranged in increasing scenario of the flood frequency growth curves (1.4. increasing LCV of Coefficients of variation and meruman computed using EVI and GEV distributions and method of moments using FWI making AFS for log-parson Type-III.		first and a high at	second a tion ake	itation o W. respe	ctively.	region ve	1	ted be	l è -	101
Coefficients of variation and skewness computed usin EV1 and GEV distributions and method of moments usin raw maximum AFS for log-Pearson Type-III.		cons are a d framman	rrangad 1 ov arowth	La Increa	41ng 85.	a apnesa 4		LSK)	6u1 pu	
Ev1 and GEV distributions and method of moments usin raw maximum APS for log-Paarson Type-III.	0	ficients	of variat	lon and	skewnee	e compute	d using	L-mome	nte for	
raw maximum APS for log-Pearson Type-III.			istribut!	ons and	method -	of moment	ce uning	normal.	Ized	
	Ner.	A BURLAUR A	PS for 10	9-PARENO						

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TABLE 1.27. Comparison of Quantiles at a Yew Balacted Stations:

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TABLE 3.28. Comparison of Quantiles of a Few Selected Stations:

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185500 1.0 26 1.0 26 1.0 26 1.0 2.0 2.0 2.0 2.0 2.0 2.0 2.0 2.0 2.0 2			-125
285300 1.0 26 0 403500 960.0 46 0 209573 3.2 10 0 297000 5.3 29 0 610830 0.1 9 0 208500 86.0 58 0 313600 1.0 1.8 0 1.0 1.0 1.4 0 1.1 1 The first and second station of and second station of a		-11 1211	Ţ
265400 1.0 26 0 403500 960.0 46 0 209573 3.2 10 0 209500 5.3 29 0 610820 6.1 9 0 209500 6.0 54 0 313600 1.0 34 0 103100 39.7 32 0 11 The first and second station of	1700	0161	;
403500 960.0 46 0 209373 3.2 10 0 297000 5.3 29 0 610830 5.3 29 0 201000 5.0 59 0 313600 1.0 18 0 103100 39.7 32 0 103100 39.8 24 0 103100 39.8 24 0 1106 6 0 0 0 0 0 0 0		2	7
403500 960.0 46 0 209573 3.2 10 0 297000 5.3 29 0 610830 6.1 9 0 208500 66.0 59 0 313600 1.0 18 0 313600 1.0 18 0 313600 1.0 18 0 10 10 18 0 10 10 0 0 10 10 10 10 10 10 10 0 10 10 10 10 0 10 0 0 0	1001	-10 1494	7
403300 940.0 44 0 209373 3.3 10 0 297000 5.3 29 0 610820 6.0 59 0 208300 66.0 59 0 113600 1.0 18 0 103100 39.7 32 0 103100 39.8 24 0 103100 39.8 24 0 103100 39.8 24 0			;
209373 3.2 10 20 29 29 29 20 29 20 20 20 20 20 20 20 20 20 20 20 20 20	67939	13 76151	1
209375 3.2 10 0 297000 5.2 29 0 610820 6.1 9 0 208500 6.0 54 0 313600 1.0 34 0 10000 89.7 32 0 103100 39.7 32 0 103100 39.8 24 0	75641	1661	ž
205375 5.2 10 0 29700 5.2 29 610830 5.3 29 206300 5.0 59 313600 1.0 18 0 113600 1.0 18 0 103100 39.7 32 0 303100 39.8 24 0 103100 19.8 24 0 114 first and second station of		3	1
297000 5.3 29 00 610830 5.3 29 0 208500 6.0 59 0 313600 1.0 18 0 1.0 18 0 303300 39.7 32 0 10 3000 39.7 32 0 11 1	797		7
297000 5.1 29 0 410810 5.1 29 0 208300 66.0 54 0 113600 1.0 14 0 113600 1.0 14 0 113600 19.7 32 0 103100 39.6 34 0 The first and second station of		-2 1055	î
297000 5.3 29 00 410830 6.1 9 0 206500 66.0 54 0 311500 1.0 18 0 610000 89.7 32 0 103100 39.4 24 0 103100 39.4 24 0 103100 39.4 24 0 11 1			1
610820 0.1 9 0 208500 6.0 54 0 313600 1.0 34 0 103100 1.0 34 0 103100 39.7 32 0 303100 39.8 24 0 11 1		-11 5114	Ŧ
<pre>\$10820 0.1 9 0 208300 0.1 9 0 208300 0.0 58 0 3113600 0.0 10 18 0 10 13 0 10 10 10 18 0 10 10 10 18 0 11 The first and second station of and a high station of and a high station of </pre>	\$233	1019 [*1
<pre>\$10830 0.1 9 0 208500 86.0 58 0 313600 1.0 18 0 103300 39.7 32 0 303300 39.8 34 0 303300 39.8 24 0 303300 39.8 24 0 1 The first and second station of and a high station of</pre>	5070	•	ł
204300 86.0 54 0 313600 86.0 54 0 610000 49.7 32 0 303100 39.8 34 0 7he first and second station of		^	7
208300 86.0 54 0 313600 1.0 18 0 610000 89.7 32 0 303300 39.4 24 0 303300 39.4 24 0 10 first and scond station of	111	-12 064	-
201300 16.0 54 0 313600 1.0 18 0 610000 19.7 32 0 303100 39.7 32 0 103100 39.8 24 0 The first and second station of	579		ł
313600 1.0 1.0 1.0 510000 85.7 7.2 0 001100 19.4 24 0 77he first and second station of and a high station of a high st	45487		-12
J13600 1.0 14 0 410000 49.7 32 0 303300 39.8 34 0 The first and second station of	46554	-4 56799	7
313600 1.0 1.0 1.0 610000 49.7 32 0 03100 39.4 3 0 77he first and second station of 1	50600		1
6 10000 89.7 33 0 0 101100 39.8 34 0 1 1 The first and second station of and a high station of	159	-10 060	
\$10000 \$9.7 32 0 303100 39.8 34 0 103100 39.8 34 0 103100 39.8 34 0 103100 39.8 34 0 103100 39.8 34 0 103100 39.8 34 0 103100 39.4 0 0 103100 39.4 0 0 103100 34.4 0 0 103100 34.4 0 0 103100 34.4 0 0		-26 940	î
 \$10000 49.7 32 0 303300 39.4 24 0 The first and second station of and a high station skev, respect 		1360	i
0 303300 39.4 34 0 The first and second station of and a high station sev. respect		-9 26653	1
0 303100 39.6 34 0 0 The first and second station of and a high station skev, respect		17 40995	5
J0JJ00 J9.4 24 0 J07he first and second station of and a high station station of	~		1
0 The first and second station of and a high station skew, respect		51	ĩ
1 The first and second station of and a high station skew, respect		31 12975	5
first and second station of a high station skew, respect	0267 -	- 935	1
a high station skew, respect		and hered hered on	101
		5	
geodean builden in Incleaning at a sector of the sector of	of the	corresponding	
ncy growth curves (1.4. in	23 6	LSK).	
Coefficients of variation and stevness computed	6u jan	L-moments for	

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** log-Pearson Type-III attlates conducted using Mater Resources Council guidelines - Builetin 178. "Diff" is the percentage difference of the guantile estimates realitive to the log-Pearson Type-III astimates.

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				Stati	on =	Q	uantilo	15 **	
luster	Station	Area	Yrs of	Valu		50 YT	Diff.	100 YE	Diff.
Region	No.	(sq mi)	Record	LCY	LSK	(cfs)	(\$)	(cfs)	(\$)
imbel/G	EV/LP3								
6	322100	323.0	22	0.24	0.21	11685	18	13069	20
•				0.24	0.21	12272	24	14073	29
				0.44	1.45	9890		10900	
	315885	0.2	9	0.31	0.41	127	-15	142	-19
	320000	•••	-	0.31	0.41	133	-11	153	-13
				0.62	2.11	150		175	
1	247100	3.3	31	0.16	0.07	13.69	45	1533	53
•	247100	2.2		0.16	0.07	1504	59	1773	77
				0.28	0.07	946		999	
	298535	0.7	. 10 .	0.51	0.55	497	-47	556	-55
	298333	0.7	. 10 .	0.51	0.55	546	-40	643	-48
				1.18	2.53	- 944		1240	
	283500	362.0	51	0.30	0.20	23649	-8	26536	-9
2	283500	797.0	21	0.30	0.20	24955	-3	28774	-1
				0.56	1.18	25600		29100	
			22	0.44	0.41	3627	-14	4064	-17
	237280	12.2	22	0.44	0.41	3821	-9	4406	-10
					2.97	4230		4920	
				0.99			25	1485	30
4	402020	3.0	10	0.19	-0.08	1325	29	1558	37
				0.19	-0.08	1368	49	1558	
				0.33	0.74	1060		7482	-18
	404900	53.8	29	0.31	0.34	6675	-14		
		•		0.31	0.34	6892	-11	7850	-14
				0.61	1.95	7760		9160	
7	610503	0.8	10	0.24	-0.06	1797	13	2760	60
				0.24	-0.06	1974	24	3827	121
•				0.40	-0.58	1590		1730	
	302500	194.0	45	0.35	0.33	21865	-18	24586	-23
				0.35	0.33	24025	-19	28401	-11
				0.71	2.17	26800		31900	
3	284300	28.6	15	0.32	0.09	8857	-16	9971	-18
				0.32	0.09	9745	-7	11540	-5
				0.56	0.18	10500		12100	
	208600	202.0	29	0.40	0.40	41858	-18	47124	-24
				0.40	0.40	46055	-10	5453 9	-12
				0.86	2.26	51300		62200	
5	415700	4.8	24	0.27	0.01	1949	- 4	2197	7
-				0.27	0.01	2177	15	2604	26
				0.47	0.26	1870		2060	
	307000	173.0	46	0.36	0.42	27952	-22	31500	29
			-	0.36	0.42	31219	-13	37336	-16
				0.77	1.88	36000		44600	

TABLE 3.29. Comparison of Quantiles of a Yew Selected Stations

The first and second station of each region were selected based on a low and a high station skew, respectively. Regions are arranged in increasing steepness of the corresponding flood frequency growth curves (i.e. increasing LCV or LSK). Coefficients of variation and skewness computed using L-moments for EV1 and GEV distributions and method of moments using normalized raw maximum AFS for log-Pearson Type-III. log-Pearson Type-III estimates computed using Water Resources Council muldilings - Building 172 Building is the normalized of ŧ

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** guidelines - Bulletin 178. "Diff" is the percentage difference of the quantile estimates realitive to the log-Pearson Type-III estimates.

c) Discriminant Analysis: In the previous section, homogeneous flood regions are identified using four different clustering schemes (referred to as Cases 1-4) using FASTCLUS clustering algorithm. Although a comparison between the cluster regions is made using important statistics of all hydrological attributes and the performance of the regionalized flood frequency growth curves, it remains to be seen as to what factors, other than the clustering variables employed, cause the fundamental differences between these regions. For instance, if cluster regions are delineated using LCV and QSP (Case 1), the five regions identified can be generally classified as low, medium and high flood regions based on the values of the response variables. Since watershed drainage area, $A_{c}^{}$, is highly correlated with one of the clustering variables, namely QSP, the differences between these clusters regions can be further explained on the basis of this or other physical attributes that control flood response of a watershed. With this in mind, a stepwise discriminant analysis is first performed in order to identify the most significant (at 5% level of significance) attribute variables which provide maximum discrimination between the flood regions. Application of this procedure to the cluster regions for all the clustering schemes gave results as summarized in Table 3.30. Results of clustering cases 1-4 and USGS regions are also included this table. The variables listed in column 3 of this table are the significant attribute variables arranged in the order of importance. The following conclusions are drawn:

 a) Although the original set of attribute variables defined at each gauge incorporated a broad range of hydrological characteristics of each watershed, the most important variables controlling flood response seem to be the geomorphic properties of the watershed

Case No.		Signif. Discrim. Variables		crimination Percent
1	LCV	DAREA, CHANSLOP	62/253	0.245
2	LCV, LSKEW	DAREA, BASLEN	65/242	0.269
3	LCV, LSKEW, LKUR	DAREA, BASLEN	72/242	0.298
4	LCV, QSP	DAREA, BASLEN	111/242	0.459
5*	LCV, LSKEW, QSP	DAREA, BASLEN	105/242	0.434
6*	LCV, LSKEW, LKUR, QSP	DAREA, BASLEN, CHANSLOP	102/242	0.421
7	MEVL, AEVL	DAREA, CHANSLOP	62/253	0.245
8	MGVL, AGVL, KGVL	DAREA, BASLEN SHAPE	79/242	0.326
9	MWKL, AWKL, BWKL, CWKL, DWKL	DAREA	159/253	0.628
10*	MEVL, AEVL, QSP	DAREA, BASLEN, SHAPE, CHANSLOP, STOR, CHANSIN, CH	•	0.471
11 [±]	MGVL, AGVL, KGVL, QSP	DAREA, BASLEN CHANSIN, CHANLEN	88/242	0.364
12	MWKL,AWKL, BWKL, CWKL, DWKL, QSP	DAREA	122/242	0.504
13	USGS REGIONS	DAREA, CHANLEN	41/242	0.169
	available for e	sical characteristic ach station, some so the discriminant an	tations are	
*	Clustering cases (referred to as	s selected in the st	tudy	

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Table 3.30. Results of Discriminant Results for all Flood Region Delineation Cases Examined in the Study ** such as its size and shape and main channel characteristics.

- b) Watershed drainage area, A_C, is the most significant attribute for discriminating between clusters for all the clustering schemes and USGS regions.
- c) All the significant attributes listed in Table 3.30 describe the physical characteristics that control the magnitude and timing of flood peak response of a watershed. For instance, the magnitude of the flood peak is proportional to the drainage area and its timing is influenced by travel paths such as the basin and main channel lengths.

The next step in discriminant analysis is to perform a classificatory analysis of gauged sites in each cluster region (for a given clustering scheme) in order to determine the percentage gauged sites correctly classified in the original cluster regions. To accomplish this the significant attribute variables are used together with the DISCRIM procedure of SAS (1985) to perform a classificatory discriminant analysis. Tables 3.31-3.34 summarize the results for all the four clustering schemes (Cases 1-4) selected in the study. The horizontal rows in these tables reflect the original cluster groupings while the vertical columns indicate the new cluster groupings into which each site is classified based upon its attributes. If all gauges are correctly classified then the row percentages of the diagonal elements in these tables will be 100%. It is obvious that such is not the case. The low percent classification in some cases indicates that the cluster regions can not be discriminated well based upon the attributes used in the analysis. An overall discriminant score is computed by summing the sites classified correctly (i.e. all sites along the diagonal). This total score divided by the total number of sites being classified gives the overall percent correct classification. This value for

each clustering scheme is shown in Column 5 of Table 3.30. Based on these results the following conclusions are drawn:

- a) With the exception of the clustering cases involving the Wakeby probability distribution parameters (these cases were dropped due poor performance in estimating flood quantiles), clustering cases 1-4 (labeled as cases 4, 5, 10 and 11 in Table 3.30) provide the best overall percent classification compared to all the other cases considered in the study. The overall percent correct classification ranges from 36.4% to 47.1%.
- b) Watershed drainage area, A_c, is the most significant discriminating variable for all the clustering schemes. The remaining variables listed in Table 3.30 are all geomorphic that are closely related to the physical aspects of the watershed.
- c) In all clustering schemes there are at least two cluster regions that have a percent classification less than 50%. This occurs with cluster regions that have considerable overlap in their hydrological characteristics.

The results of discriminant analysis for the seven USGS flood regions using all the gauged site data as illustrated in Table 3.35. An examination of the diagonal elements of this table clearly indicates that these regions can not be discriminated between each other easily. In other words, the classification of gauged sites into a region based upon the attribute variables (referred to as discriminating power) can not be achieved with a high degree of certainty. The average discrimination is only 16.9% when compared to a maximum of 47.1% achieved using cluster analysis in conjunction with EV1 parameters and QSP as the clustering variables (refer to Case 10 in Table 3.30). This further supports the observation that each of the USGS flood regions

has a mixed composition of watersheds with differing hydrological characteristics. Hence, these regions are not very homogeneous with respect to the characteristics describing flood response. A similar observation was made by Wiltshire (1986) who states that flood regions delineated in a rather arbitrary manner and arranged to coincide with geographical areas are likely to contain drainage basins with a diversity of geomorphology whose flood frequency characteristics may not be comparable. He further states that in this situation a regional average frequency curve will be poorly defined. In contrast, the cluster regions are not only homogeneous with respect to the flood response characteristics (response or clustering variables) but lend themselves to a higher level of discrimination by variables that affect these characteristics (attribute variables). Α comparison of the significant variables in the discriminant analysis indicates that for both cluster and USGS regions the geomorphic variables provide good discrimination with contributing drainage area being the most important (refer to Table 3.30). The remaining variables describe the watershed and main channel dimensions.

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From Cluster		•	-	4	5	Marka 1
CIUSCEL	• •	2	2	4	5	Total
1	49	6	20	11	3	89
	55.1	6.7	22.5	12.4	3.4	100.0
2	2	8	0	2	2	14
	14.3	57.1	0 0.0	14.3	14.3	100.0
3	38				5 5.6	90
	42.2	3.3	34.4	14.4	5.6	100.0
4	o	6	0	21	0	27
	0.0	22.2	0.0	77.8	0.0	100.0
5	5	7		7	2	22
	22.7	31.8	4.6	31.8	9.1	100.0
		30				242
Percent	38.8	12.4	21.5	22.3	5.0	100.0
Priors	0.20	0.20	0.20	0.20	0.20	
Overall %		+ clas		tion -		2 - 165

Table 3.31. Classificatory Discriminant Analysis of Cluster Regions Formed Using Clustering Variables LCV and QSP *

Number and Percentage of Observations

basin length, B₁, basin shape, B_s, and main channel^c, slope, S_c.

Table 3.32. Classificatory Discriminant Analysis of Cluster Regions Formed Using Clustering Variables EV1 Parameters and QSP *

From Cluste	r 1	2	3	4	5	Total
1	5 12.8				11 28.2	39 100.0
2	2 22.2		0 0.0		1 11.1	9 100.0
3	4 4.4				12 13.2	91 100.0
4	3 3.9	3 3.9			12 15.6	77 100.0
5	2 7.7		0 0.0		19 73.1	26 100.0
otal cent	16 6.6		103 42.6		55 22.7	242 100.0
iors	0.20	0.20	0.20	0.20	0.20	

Number and Percentage of Observations Classified into Cluster Region

Significant variables at S_1 level: drainage area, A_c, basin length, B₁, basin shape, B_s, main channel slope, S_c, main channel sinuousity, S_s, and basin storage, STOR.

From	•	-	-		_		
CIUSC	er 1	2	د	4	5	Total	
1	21	19	17	3	19	79 100.0	
	26.6	24.1	21.5	3.8	24.1	100.0	
2	2	4	1	0	7	14 100.0	
	14.3	28.6	7.1	0.0	50.0	100.0	
3	15	7	48	1	4	75 100.0	
	20.0	9.3	64.0	1.3	5.3	100.0	
4	12	9	8	o	11	40 100.0	
	30.0	22.5	20.0	0.0	27.5	100.0	
5	0	2	0	٥	32	34	
	0.0	16.9	30.6	1.7	30.2	34 100.0	
Total	50	41	74	4	73	242	
rcent	20.7	16.9	30.6	1.6	30.2	100.0	
riors	0.20	0:20	0.20	0.20	0.20		
Correct	classi	ficati	on = 1	05/242	= 431	**********	

Table 3.33. Classificatory Discriminant Analysis of Cluster Regions Formed Using Clustering Variables LCV, LSK and QSP *

Number and Percentage of Observations Classified into Cluster Region

Table 3.34. Classificatory Discriminant Analysis of Cluster Regions Formed Using Clustering Variables GEV Parameters and QSP \pm

1 2	47 58.0	0					
	58.0		13	8	8	5	81
2		0.0	16.1	9.9	9.9	6.2	100.0
	1	4	2	.11	0	0	18
	5.6	22.2	11.1	61.1	0.0	0.0	100.0
3	11	5	7	12	2	2	39
	28.2	12.8	18.0	30.8	5.1	5.1	100.0
4	0	2	2	10	0	0	14
	0.0	14.3	14.3	71.4	0.0	0.0	100.0
5	18	7		6	19	3	64
	28.1	10.9	17.2	9.4	29.7	4.7	100.0
6	4	5	5				26
	15.4	19.2	19.2	30.8	11.5	4.6	100.0
Total	81	23	40		32	11	242
Percent	33.5	9.5	16.5	22.7	13.2	4.6	100.0
Priors 0	.167	0.167 0	.167 ().167	0.167 0	.167	
<pre>% Correct</pre>	class:	ificati	on =	88/24	2 = 36%		· · · · · · · · · · · · · · · · · · ·

Number and Percentage of Observations Classified into Cluster Region

From		-			-	-	_	
Cluster	1	2	3	4	5	6	. 7	Total
1	22	0	0	ı	0	5	2 6.7	81
	73.3	0.0	16.1	1 9.9	9.9	16.7	6.7	100.0
2	29	0	2	13	0	18	6	. 68
	42.7	0.0	2.9	19.1	0.0	26.5	8.8.	100.0
3	8	0	2	5	0	6		26
	30.8	0.0	7.7	19.2	0.0	23.1	19.2	100.0
4	-10	2	0	5	0	2	1	20
	50.0	10.0	0.0	25.0	0.0	10.0	5.0	100.0
5	17	0	3	. 4	0	. 6	5	35
	48.6	0.0	8.6	11.4	0.0	17.1	14.3	100.0
6	12	0	2	2	1	9	2	28
	42.7	0.0	7.1	7.1	3.6	32.1	7.1	100.0
7	20	0		0	. 1	10	3	35
·	57.1	0.0	2.3	0.0	2.3	28.6	8.6	100.0
otal	118					56		242
ercent	48.8	0.8	. 4.1	12.4	0.8	23.1	9.9	100.0
riors	0.142	0.142	0.142	0.142	0.142	0.142	0.142	
Correc	t clas	ssifica	tion =	= 41/2	242 = 1	 17%		

Table 3.35. Classificatory Discriminant Analysis of USGS Regions *

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The ultimate objective or purpose d) Regression Analysis: of delineating distinct flood regions is to develop regionalized relationships for predicting the flood response (at selected frequency levels) at both gauged and ungauged sites. For gauged sites, such a regionalized relationship can be used together with at-site information for estimating flood levels (Choquette 1988). In contrast, while using cluster analysis, ungauged sites must first be classified into a flood region based on significant physical attributes of the watershed affecting flood response prior to using a regionalized equation. For the method of residuals, this classification is relatively straight forward since an ungauged site is univocally assigned to the geographic region in which it lies.

Overall regression results, pertaining to the equations developed for predicting the 20, 50 and 100 year flood levels within the cluster regions (for all clustering cases examined in this study), are shown in Table 3.36. Cases 1-4 (marked by an asterisk in this table) have the lowest weighted standard error when compared to the remaining cases. For these four cases and USGS flood regions, detailed regionalized regression equations for the EV1 and GEV models are given in Tables 3.37-3.46.

Table 3.47 gives similar equations for the 50 and 100 year flood levels (20 year flood quantile regression equations are not available) and are developed for the USGS method of residuals flood regions using log-Pearson Type-III distribution (Choquette, 1988). This table is provided for the purpose of comparing the performance of the log-Pearson Type-III, EV1 and GEV flood frequency models developed for the seven USGS flood regions. It must be emphasized that the flood levels used in developing these regression equations are estimated from a log-Pearson Type-III flood frequency distribution using a weighted skewness (based on station and a map skew).

Case No.	Cluster Variables	No. of Regions	Weighted Standard Error (Standard Error Range** (%)
1	LCV	6	44.4	32.8 - 53.0
2	LCV, LSKEW	6	44.9	39.8 - 56.5
3	LCV, LSKEW, LKUR	6	45.2	39.2 - 53.1
4*	LCV, QSP	5	36.9	19.3 - 46.6
5*	LCV, LSKEW, QSP	5	36.6	27.0 - 51.0
6	LCV, LSKEW, LKUR, QSP	5	41.1	24.9 - 54.9
7	MEVL, AEVL	6	43.8	32.2 - 53.2
8	MGVL, AGVL, KGVL	5	44.5	39.0 - 56.1
9	MWKL, AWKL, BWKL CWKL, DWKL	, 2	45.8	44.0 - 54.1
10*	MEVL, AEVL, QSP	5	39.2	20.1 - 54.9
11*	MGVL, AGVL, KGVL, QSP	6	39.1	23.4 - 52.2
12	MWKL, AWKL, BWKL CWKL, DWKL, QSP	, 6	45.7	43.5 - 115.6
13	USGS REGIONS	7	35.0	19.7 - 38.6
() ** Ba	ndicates cases whi referred to as Cas ased on standard o agion	ses 1-4)		-

TABLE 3.36. Regression Results for all Flood Region Delineation Cases Examined in the Study

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An examination of these tables suggests that the standard errors associated with the regression equations are, in general, comparable between the cluster and USGS regions when EV1 and GEV models are used in the regionalization. However, the standard errors are slightly higher when using the log-Pearson Type-III distribution (compare Table 3.46 and 3.47). Hence, even for the USGS flood regions (as delineated using method of residuals), it appears that more accurate regression equations can be developed by using either EV1 or GEV regionalized flood frequency models.

The independent variables and their exponents do not change for a particular flood frequency model within a flood region for different return period, T. This is not surprising since the regionalized quantile levels (normalized values) used in estimating the flood levels, $Q_{m},$ are scalar multiples of each other. In other words, the 100-year flood quantile can be obtained from the 10-year flood quantile by multiplying the latter with a constant. Hence, the correlation of flood quantiles with the independent variables remains the same from one flood level to the next. Consequently, the exponent term in the regression equations (slope term in the log-relationship) remains unaffected. The effects of scale are absorbed in the intercept term. A similar reasoning applies when comparing the regression equations (for a given flood quantile, ${\tt Q}_{m},$ and flood region) between the EV1 and GEV flood frequency models. Once again, the independent variables and their exponents continue to be identical within a flood region when the return period, T is changed.

For both cluster and USGS regions, the geomorphic variables such as the watershed drainage area, A_c , main channel slope, S_c , and sinuousity, S_s , basin shape, B_s , are the most significant variables. For some cluster regions the exponent of the independent variable (drainage area, A_c) is greater than or equal to 1.0, indicating greater

variability in the estimate of the flood quantile as the drainage area increases (true for watersheds greater than 1.0 sq mi). These cluster regions, as compared to other regions have predominantly small watersheds. Furthermore, gauged sites within these cluster regions also have, in general, short flood records (less than 10 years). Thus, a possible explanation for the larger exponent of the drainage area variable, A_c , in the regression equations may be due the fact that small watersheds experience greater variability in their flood response as opposed to larger watersheds due to their inability to dampen temporal effects of rainfall.

In applying the regionalized regressions equations developed for cluster regions (Cases 1-4) for ungauged sites (these sites do not have their flood characteristics defined), one must first assign these sites to a particular cluster region based solely on the physical attributes of the watersheds. Results of classificatory discriminant. analysis (refer to Tables 3.31-3.34) of the gauged sites, based on their physical attributes only (i.e. treating them as ungauged sites), show the assignment of watersheds is not with complete certainty. For instance when clustering with LCV and QSP (Case 1), 49 of the 89 sites originally assigned to cluster region 1 are re-assigned to this region while the remaining sites are assigned to the cluster region 2, 3, 4 The posterior probabilities of these assignments are or 5. given in the second row of Table 3.31 for each cluster region. Consequently, in using the regionalized regression equations shown in Table 3.37 for predicting the flood levels at ungauged sites, one must use a weighted predicted flood level as developed from all the regionalized regression equations associated with the cluster regions to which the site is assigned. The weighting can be accomplished using the posterior probabilities of being assigned to each cluster region. Thus, the standard errors of prediction must also be based on the regionalized

regressions used. For each cluster region, a weighted standard error may be computed using the following equation:

$$m
 e_{j} = \sum_{ji}^{m} e_{ji} for j = 1, 2...m ... (3.1)
 i = 1$$

where,

ej = percent standard error at a site in cluster j, eji = percent standard error at a site in cluster j if it was classified into cluster i, pji = posterior probability of a site in cluster j being classified into cluster i, and m = number of cluster regions (equal to 5 or 6 in

the present study).

For clustering Cases 1-4, values of e_{ji} can be obtained from the standard errors shown in Column 3 of Table 3.37-3.45 for each cluster region i and the posterior probabilities, p_{ji} , can be obtained from the rows of Tables 3.31-3.34. Based on all the gauged sites that are classified into the cluster regions, a weighted standard error can be computed for each cluster region using Eq. 3.1 above. For USGS regions, the problem of misclassification does not arise since ungauged sites are assigned to a region on the basis of their location in space.

Cluste Region		% Standard Error	R ²	No. of Sites.
1	Q20 = 546 Ac 1.061 L -0.552	38.6	0.95	88
2	$Q_{20} = 967 \lambda_{c}^{0.716} s_{c}^{-0.168}$	19.3	0.95	15
3	$Q_{20} = 382 \lambda_c^{0.777} s_c^{0.144}$	38.1	0.94	92
4	Q ₂₀ = 803 A _c ^{0.960}	29.0	0.93	29
5	$Q_{20} = 657 \lambda_c^{0.682} s_c^{-0.346}$	46.6	0.93	24
1	$Q_{50} = 659 \lambda_a^{1.061} L_a^{-0.552}$	38.6	0.95	88
2	$Q_{50} = 1183 A_{c}^{0.716} S_{c}^{-0.16}$	8 19.3	0.95	15
3	$Q_{50} = 449 \lambda_c^{0.777} s_c^{0.144}$	38.1	0.94	92
4	Q ₅₀ - 958 д _а ^{0,960}	29.0	0.93	29
5	$Q_{50} = 812 \lambda_{c}^{0.682} s_{c}^{-0.346}$	46.6	0.93	24
	$Q_{100} = 742 \lambda_{c}^{1.061} L_{c}^{-0.552}$	-	0.95	88
2	$Q_{100} = 1345 \lambda_c^{0.716} s_c^{-0.163}$	8 19.3	0.95	15
3	$Q_{100} = 499 \lambda_c^{0.777} s_c^{0.144}$	38.1	0.94	92
4	$Q_{100} = 1077 \ M_{c}^{0.960}$	29.0	0.93	29
5	$Q_{100} = 927 \lambda_c^{0.682} s_c^{-0.346}$	46.6	0.93	24

TABLE 3.37. Regression Models for Estimating the Expected EV1 Flood Quantiles for Various Return Periods: Clustering ' on LCV and QSP (Case 1)

Cluste Regior		t Standard Error	R ²	No. of Sites.
1	Q20 - 550 AC 0.758	54.9	0.90	42
2	$Q_{20} = 516 L_{C}^{1.596}$	20.1	0.90	8
3	$Q_{20} = 495 \lambda_c^{0.975} L_c^{-0.402}$	37.2	0.95	90
4	$Q_{20} = 397 \lambda_c^{0.777} s_c^{0.169}$	38.6	0.93	78
5	$Q_{20} = 781 \lambda_{C}^{0.980}$	29.5	0.93	29
1	$Q_{50} = 676 \lambda_{C}^{0.758}$	54.9	0.90	42
2	$Q_{50} = 629 L_{C}^{1.596}$	20.1	.0.90	8
3	$Q_{50} = 592 \lambda_{c}^{0.975} L_{c}^{-0.402}$	37.2	0.95	90
4	Q ₅₀ - 466 $\lambda_{2}^{0.777}$ s ₂ ^{0.169}	38.6	0.93	78
5	$Q_{50} = 928 \lambda_{G}^{0.980}$	29.5	0.93	29
1	$Q_{100} = 773 \lambda_G^{0.758}$	54.9	0.90	42
2	$Q_{100} = 711 L_{c}^{1.596}$	20.1	0.90	8
3	$Q_{100} = 669 \lambda_c^{0.975} L_c^{-0.402}$	37.2	0.95	90
4	$Q_{100} = 517 \lambda_c^{0.777} s_c^{0.169}$	38.6	0.93	78
	0.980			

29.5

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0.93 29

TABLE 3.38. Regression Models for Estimating the Expected EVI Flood Quantiles for Various Return Periods: Clustering on EV1 Parameters and QSP (Case 2)

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 $Q_{100} = 1043 \ A_{C}^{0.980}$

TABLE 3.39: Regression Models for Estimating the Expected EV1 Flood Quantiles for Various Return Pariods: Clustering on LCV, LSK and QSP (Case 3)

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Cluster Region		Regression Equation	<pre>\$ Standard Error</pre>	'R ²	No. of Sites,
1	Q ₂₀	- 574 A 1.069 L -0.605 S -0.	119 40.4	0.95	78
2	Q	- 659 A_0.776	51.0	0.90	16
3	0	= 395 $\lambda_{0.821}^{0.821}$ 5 $\lambda_{0.227}^{0.227}$ B $^{-0.1}$	94 32.7	0.91	74
4	Q ₂₀	= 465 $\lambda_{c}^{0.793} s_{s}^{-0.911}$	40.1	0.95	39
5	Q ₂₀	- 887 А. 0.887 С	27.0	0.91	37 [.]
1	Q ₅₀		¹¹⁹ 40.4	0.95	78
2	Q ₅₀	= 816 A 0.776	51.0	0.90	16
Э	950	$= 467 \lambda_{0.621}^{0.621} \text{ s}_{0.227}^{0.227} \text{ s}_{-0.1}^{-0.1}$.94 32.7	0.91	74
4	Q ₅₀	- 544 A 0.793 8 -0.911	40.1	0.95	39
5		- 1072 A 0.887	.27.0	0,91	37,
1	Q ₁₀₀		119 40.4	0.95	78
2	Q ₁₀₀	= 936 A _c ^{0.776}	51.0	0.90	16
3	0	$= 524 \lambda_{0.021}^{0.021} s_{0.227}^{0.1} B_{-0.1}^{-0.1}$.94 32.7	0.91	74
4	Q ₁₀₀	= $602 \lambda_{c}^{0.793} s_{-0.911}^{-0.911}$	40.1	0.95	39
5	Q100	= 1207 A 0.887	27.0	0.91	37

Cluste Region		t Standard Error	R ²	No. of Sites.
1	$Q_{20} = 480 \ A_{c}^{0.963} \ L_{c}^{-0.396}$	38.1	0.94	80
2	$Q_{20} = 893 \lambda_c^{0.850}$	- 23.4	0.94	20
3	$Q_{20} = 661 \lambda_{c}^{1.225} L_{c}^{-0.809}$	52.2	0.93	38
4	Q20 = 904 Ac 1.002	31.2	0.91	14
5	Q20 = 566 AC0.962 L -0.456	39.1	0.95	63
6	$Q_{20} = 688 \lambda_c^{0.643} s_c^{-0.206}$	42.0	0.95	27
1	$Q_{50} = 584 \lambda_{c}^{0.963} L_{c}^{-0.396}$	38.1	0,94	80
2	Q ₅₀ = 1080 λ _c ^{0.850}	23.4	0.94	20
3	$Q_{50} = 803 \lambda_c^{1.225} L_c^{-0.809}$	52.2	0.93	38
4	$Q_{50} = 1076 \lambda_{c}^{1.002}$	31.2	0.91	14
5	$Q_{50} = 666 \lambda_{c}^{0.962} L_{c}^{-0.456}$	38.1	0.95	63
6	$Q_{50} = 847 \lambda_{c}^{0.643} s_{c}^{-0.206}$	42.0	0.95	27
1	Q ₁₀₀ = 640 A _a ^{0.963} L _a ^{-0.396}	38.1	0.94	80
2	$Q_{100} = 1217 \lambda_{c}^{0.850}$	23.4	0.94	20
3	$Q_{100} = 910 \lambda_c^{1.225} L_c^{-0.809}$	52.2	0.93	38.
4	$Q_{100} = 1206 \lambda_c^{1.002}$	31.2	0.91	14
5	$Q_{100} = 741 \lambda_{C}^{0.962} L_{C}^{-0.456}$	30.1	0.95	63
6	$Q_{100} = 967 \lambda_{c}^{0.643} s_{c}^{-0.206}$	42.0	0.95	27
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TABLE 3.40. Regression Models for Estimating the Expected EV1 Flood Quantiles for Various Return Periods: Clustering on GEV Parameters and QSP (Case 4)

	ster ions		Regress Equati		t s	tandard Error	r ²	No. of Sites.
1	Q ₂₀	= 716	λ_ ^{0.963}	L _c ^{-0.396} s	20.196	38.4	0.95	29
2	Q ₂₀	= 341	A_0.736			39.9	0.96	67
3	Q ₂₀	- 520	λ.0.744	sc ^{0.029} Bs	-0.070	19.7	0.99	25
4	Q ₂₀	= 289	λ _c ^{0.842}	s0.517		27.1	0.99	19
5	Q ₂₀	= 634	A_0.720			33.2	0.97	37
6	Q ₂₀	= 623	λ _C	s_ ^{-0.277}		38.6	0.96	27
7	Q ₂₀	= 818	Ac 0.587	•		36.9	0.95	37
1	Q ₅₀	- 852		L0.396 S	20.196	38.4	0.95	29
2	0 ₅₀	- 408				39.9	0.96	67
з	Q ₅₀	= 622	<u> </u>	s _c ^{0.029} B	-0.070	19.7	0.99	25
4	Q ₅₀	= 344	a.0.842	s0.517 s		27.1	0.99	19
5	Q ₅₀	= 763	a.0.720			33.2	0.97	37
6	0 ₅₀	- 739	5	s0.277		38.6	0.96	27
7	Q ₅₀	= 982	λ.587		-	36.9	0.95	37
1	Q ₁₀₀	- 954	م 0.963 م	L _c -0.396 s	5_0.196	·38.4	0.95	29
2	0 ₁₀₀	= 457	ا م ^{0.736}			39.9	0.96	67
3	Q ₁₀₀	= 702	د م. ^{0.744}	sc ^{0.029} B	-0.070 B	19.7	0.99	25
4	Q ₁₀₀	= 385	5 A 0.842	s_ ^{-0.517}		27.1	0.99	19
5	0	- 860	ג 0-720 א			33.2	0.97	37
6	Qioo	- 830) A _C 0.624	s.=0.2//		38.6	0.96	27
7	Q ₁₀₀	= 110	0.58	7		36.9	0.95	37

TABLE 3.41. Regression Models for Estimating the Expected EVI Flood Quantiles for Various Return Periods for USGS Regions

Cluster Region		Regression Equation	t Stan Erre	dard R ² or	No. of Sites.
1	Q ₂₀	- 412 Å 1.061 L -0.552	38.6	0.95	68
	0	= 1013 A 0.716 s -0.168	19,3	0.95	15
	Q ₂₀	= 415 $A_c^{0.777} S_c^{0.144}$	38.1	0.94	92
4		= 816 Å	29.0	0.93	29
5	Q ₂₀	= 696 A 0.682 S -0.346	46.6	0.93	24
1	Q ₅₀	= 530 $\lambda_{c}^{1.061} L_{c}^{-0.552}$	38.6	·0 .95	88
2	Q ₅₀	= 1358 A 0.716 S -0.168	19.3	0.95	15
Э	Q ₅₀	= 488 A 0.777 S 0.144	38.1	0.94	92
4	Q ₅₀	= 987 A 0.960	29.0	0.93	29
5	Q ₅₀	= 1018 A _G ^{0.682} S _C ^{-0.346}	46.6	0.93	24
1 (2 ₁₀₀	= 631 $\lambda_{c}^{1.061} L_{c}^{-0.552}$	38.6	0.95	88
2 (2 ₁₀₀	= 1661 $\lambda_{c}^{0.716} s_{c}^{-0.168}$	19.3	0.95	15
3 (2100	= 544 A _c 0.777 S _c 0.144	38.1	0.94	92
4 0	2 ₁₀₀	= 1121 $\lambda_{0.960}^{0.960}$	29.0	0.93	29
		= 1332 $\lambda_{g}^{0.682} s_{g}^{-0.346}$	46.6	0.93	24

TABLE 3.42: Regression Models for Estimating the Expected GEV Flood Quantiles for Various Return Periods: Clustering on LCV and QSP (Case 1)

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Cluste Region		Regression Equation	t Standard Error	R ²	No. of Sites,
1	Q ₂₀	- 582 A C.758	54.9	0.90	42
2		$= 537 L_{c}^{1.596}$	20.1	0.90	8
C	Q ₂₀	= 509 $\lambda_{c}^{0.975} L_{c}^{-0.402}$	37.2	0.95	90
4	Q ₂₀	= 395 $\lambda_c^{0.777} s_a^{0.169}$	38.6	0.93	78
5	Q ₂₀	- 789 A C.980	29.5	0.93	29
1	Q ₅₀	= 824 A 0.758	54.9	0.90	42
2	Q ₅₀	~ 704 L_1.596	20.1	0.90	8
3	Q ₅₀	= 649 $\lambda_{c}^{0.975} L_{c}^{-0.402}$	37.2	0.95	90
4	Q ₅₀	= 464 Å C	38.6	0.93	78
5	0 ₅₀	= 956 A _C 0.980	29.5	0.93	29
1 (2 ₁₀₀	= 1053 $A_{c}^{0.758}$	54.9	0.90	42
2 (2 ₁₀₀	= 847 $L_{c}^{1.596}$	20.1	0.90	8
з (2100	= 766 $\lambda_{c}^{0.975} L_{c}^{-0.402}$	37.2	0.95	90
4 (2100	= 512 $\lambda_{c}^{0.777}$ s $c_{c}^{0.169}$	38.6	0.93	78
5 <u>(</u>	2100	= 1086 A C.980	29,5	0.93	29

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	star ion No.	Régression Equation	Standard , Error	R ²	No. of Sites.
1		601 Ac 1.069 L -0.605 S	0.119 40.4	0.95	78
2	Q ₂₀ -	694 A C	51.0	0.90	16
э	Q ₂₀ =	399 A_0.821 S_0.227 B0	.194 32.7	0.91	74
4	Q ₂₀ =	444 A 0.793 S -0.911	40.1	0.95	39
5	Q ₂₀ =	912 A 0.887	27.0	0.91	37
1	0 ₅₀ -	802 A 1.069 L ~0.605 S -	0.119 40.4	0.95	78
2	Q ₅₀ =	1058 A 0.776	51.0	0.90	16
3	Q ₅₀ -	482 A 0.821 S 0.227 B -0	.194 32.7	0.91	74
4		492 A 0.793 5 -0.911	40.1	0.95	39
5	0 ₅₀ -	1148 A 0.887	27.0	0.91	37
1	Q ₁₀₀ -	979 A _c ^{1.069} L _c ^{-0.605} S _c ⁻	-0.119 40.4	0.95	78
2	Q =	1462 A 0.776	51.0	0.90	16
C	Q =	544 A_0.821 S_0.227 B0	.194 32.7	0.91	74
4	Q ₁₀₀ =	523 A _G ^{0.793} S ₂ ^{-0.911}	40.1	0.95	. 39
5	Q ₁₀₀ =	1345 A 0.887	27.0	0.91	37

TABLE 3.44. Regression Models for Estimating the Expected GEV Flood Quantiles for Various Raturn Pariods; Clustering on LCV, LSK and QSP (Case 3)

TABLE 3.45.	Regression	Nodels :	for Est:	imating ti	he Expected
GEV Flood Qu	antiles for	Various	Return	Periods:	Clustering
on GEV Param	eters and Q	SP (Case	- 4)		

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Cluste Region		Regression Equation	\$ Standard Error	R ²	No. of Sites.
1	Q ₂₀ -	497 A _a ^{0.963} L _a ^{-0.396}	38.1	0.94	60
2	0	935 A 0.850	23.4	0.94	20
່ເ	Q ₂₀ -	679 λ 1.225 L -0.009	52.2	0.93	38
4	Q ₂₀ =	871 Ac 1.002	31.2	0.91	14
5	Q ₂₀ =	566 A 0.962 L -0.456	38.1	0.95	63
6	Q ₂₀ -	724 A 0.643 8 -0.206	42.0	0.95	27
1	Q ₅₀ -	- 635 A 0.963 L -0.396	38.1	0.94	80
2	Q ₅₀ -	- 1246 Å 0.850	23.4	0.94	20
3	Q50 •	- 857 λ 1.225 L -0.809	52.2	0.93	38
4	Q ₅₀ •	985 A 1.002	31.2	0.91	. 14
5	Q ₅₀ *	= 666 $\lambda_{c}^{0.962} L_{c}^{-0.456}$	38.1	0.95	63
6	0 ₅₀ -	= 1093 A _C ^{0.643} S _C ^{-0.206}	42.0	0.95	27
1	Q ₁₀₀ ·	- 757 A _c ^{0,963} L _c ^{-0,396}	38.1	0.94	80
2	Q100 '	- 1525 λ _c ^{0,850}	23.4	0.94	20
3	Q100 '	- 1002 Å 1.225 L -0.809	52.2	0.93	38
4	Q. 00 .	- 1062 A_ ^{1.002}	31.2	0.91	14
5	0	= 741 Å_0.902 L_0.450	38.1	0.95	; 63
-6	Q ₁₀₀	- 1476 Å C -0.206	42.0	0.95	27

Reg: No.	ion			Regres: Equat:			tandard Error	R ²	No. of Sites.
1	Q ₂₀	- 7	42	A_0.963	L _c ^{-0.396} s	0.196	38.4	0.95	29
2	2 ₂₀	- 3	48	λ _c ^{0.736}			39.9	0.96	67
3	Q ₂₀	- 6	84	λ. ^{0.744}	sc ^{0.029} Bs	-0.070	19.7	0.99	25
4	Q ₂₀	- 2	92	۸_ ^{0.842}	S0.517		27.1	0.99	19
5	Q ₂₀	- 6	58	A_0.720			33.2	0.97	37
6	9 ₂₀	- 6	i36	λ. ^{0.624}	s_ ^{-0.277}		38.6	0.96	27
7	Q ₂₀	= 8	46	λ _G 0.587			36.9	0.95	37
1	Q ₅₀	- 9	36	۸ ₀ ,963	L _c ^{-0.396} s	0.196 c	38.4	0.95	29
2	Q ₅₀	-	30	λ _G 0.736			39.9	0.96	67
3	0 ₅₀	- (84	λ _C ^{0.744}	sc ^{0.029} Bs	-0.070	19.7	0.99	25
4	Q ₅₀	= 3	54	λ _c ^{0.842}	s ^{-0.517}		27.1	0.99	19
5	Q ₅₀	- 8	51	λ _G 0.720			33.2	0.97	37
6	Q ₅₀	= 7	78	λ _c ^{0.624}	s ^{-0.277}		38.6	0.96	27
7	Q ₅₀	= 1	1077	, Y 0.23.	7 .		36.9	0.95	37
1	Q ₁₀₀	= 1	1104	1 Ac 0.96	3 _L -0.396	sc ^{0.196}	38.4	0.95	29
2	Q ₁₀₀	- 4	95	A_0.736			39.9	0.96	67
3	Q ₁₀₀	- 8	12	λ_ ^{0.744}	Sc 0.029 B	-0.070	19.7	0.99	25
4	Q ₁₀₀	= 4	104	A_0.842	s0.517		27.1	0.99	19
5	Q ₁₀₀	= 1	1020	کر انگر ا	0		33.2	0.97	37
6	Q ₁₀₀	- 8	91	λ _C ^{0.624}	s0.277		38.6	0.96	27
7	Q ₁₀₀	= 1	1273	, ^{0,58}	7		36.9	0.95	37

TABLE 3.46. Regression Models for Estimating the Expected GEV Flood Quantiles for Various Return Periods for USGS Regions

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Region No.	Regression Equation	Standard Error %	
1	$Q50 = 56 A_c^{0.959} S_c^{0.617}$	44.7	33
2	$Q50 = 670 A_{c}^{0.777} B_{s}^{-0.356} S_{s}^{-0.803}$	33.9	77
3	$Q50 = 849 A_c^{0.714} S_s^{-0.392}$	23.4	26
4	$Q50 = 363 A_c^{0.780}$	26.7	20
5	$Q50 = 940 A_{c}^{0.690}$	48.5	40
6 ·	$Q50 = 74 \lambda_c^{0.873} S_c^{0.520}$	36.2	32
7	$Q50 = 1530 A_{c}^{0.639} B_{s}^{-0.472} S_{s}^{-0.579}$	37.6	38
1	$Q100 = 51 A_c^{0.978} S_c^{0.669}$	47.8	33
2	$Q100 = 798 A_c^{0.777} B_s^{-0.373} S_s^{-0.862}$	35.1	77
3	$Q100 = 1030 A_c^{0.711} S_s^{-0.447}$	24.6	26
4	$Q100 = 420 A_{C}^{0.775}$	26.7	20
5	$Q100 = 1100 A_c^{0.689}$	52.3	40
б	$Q100 = 76 A_c^{0.882} S_c^{0.545}$	38.1	32
7	$Q100 = 1710 A_c^{0.639} B_s^{-0.466} S_s^{-0.528}$	39.4	38

TABLE 3.47. Regression Models for Estimating the Expected log-Pearson Type III Flood Quantiles for Various Return Periods for USGS Regions (Choquette, 1988) **

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** Flood Regions delineated using Method of Residuals using WRC Bulletin 17-B guidelines with a gauged site and regionalized weighted skew.

CHAPTER 4

CONCLUSIONS

Based upon the FASTCLUS algorithm, cluster analysis is used to identify distinct flood regions for the State of Important statistical properties of the annual Kentucky. maximum flood (AMF) series and other watershed hydrologic data from 253 gauged sites in the State of Kentucky are used in the analysis. Clustering variables used in the study are the L-moments, namely the coefficients of variation, LCV, and skewness, LSK of normalized maximum annual flood series, the parameters of the EV1 and GEV probability distributions, and the specific mean annual flood, QSP, based on the raw maximum annual flood series. All clustering variables are further standardized prior to clustering to suppress effects of scale. A comparison of the regions delineated under the two approaches, namely, cluster analysis and method of residuals, is then carried out using the following steps: a) direct comparison of gauged stations assigned to each region; b) comparison of mean, median and range (difference between the maximum and minimum values) of distributional characteristics of all hydrological variables (response and attribute); c) performance of regionalized EV1 and GEV flood frequency models; d) results of discriminant analysis; and e) results of regression analysis relating regionalized estimates flood quantiles of various return periods, Q_{m_i} , to watershed physical and climatic characteristics (referred to as the attribute variables). The following conclusions are made in this study:

- While the USGS method of residuals regions are or at least made to coincide with geographic or hydrologic boundaries, cluster regions do not. A comparison of actual gauged sites shows considerable difference. Cluster regions differentiate characteristics that control the underlying probability law of flood response, whereas the method of residuals does not address this issue directly.
- 2. For cluster regions the shape of the regionalized flood frequency growth curve depends on the clustering variables and the underlying probability distribution used. For EV1 distribution these growth curves are linear with normalized discharge ratio ranging from 0.0-5.0. In contrast, the GEV distribution growth curves become increasingly non-linear as the coefficients of variation and skew increase. The normalized discharge levels in this case range from 0.0-10.0. For the USGS regions the growth curves practically plot on one another indicating homogeneity of flood response between flood regions.
- 3. The regionalized EV1 and GEV flood frequency growth curves show more differences between cluster regions than between the seven USGS regions. This suggests that the cluster regions delineated for Cases 1-4 are homogeneous within themselves but distinct from each other in terms of their flood response when compared to the USGS regions. This property is essential for deriving maximum benefit from any flood regionalization effort.
- 4. An examination of statistical trends, like the mean and median, of important hydrologic variables (both response and attribute), such as the watershed area, A_c, indicates that the cluster regions have lower variability within each of the with respect to these parameters than the USGS regions. Cluster regions are

generally grouped into low, medium or large watershed drainage areas and flood response areas (as measured by the specific mean annual flood, QSP, and the mean annual flood). In contrast, the USGS flood regions have a mixed population within each of the seven regions, thereby giving similar values across regions. An exception to this are the trends observed for the basin shape, B_s, and channel sinuousity, S_s. These variables, by virtue of their definition, involving ratios of similar magnitude either small or large, show similar variation between regions for both cluster and USGS regions.

- 5. The performance of these regionalized EV1 and GEV flood frequency models, in terms of regional average bias (computed by taking the difference between simulated and historical estimates of flood quantiles) and RMSE (computed by taking the square of the bias) are comparable for cluster and USGS regions. For both models and all flood regions (cluster and USGS), the bias changes from positive to negative as the return period increases (i.e. the flood frequency growth curve becomes steeper) indicating an underestimation of flood quantiles. This trend in the bias is partly due to the condition of separation commonly found in Monte Carlo simulated flood data. This condition of separation causes simulated flow sequences to have less variability (with the separation increasing with return period) than the historical flood records resulting in an underestimation of flood quantiles when using simulated flows at a gauged site.
- 6. The absolute value of the regional average bias is less than 15% for all flood regions and flood frequency models when predicting flood quantiles having a return period less than 100 years. This indicates a high level of accuracy in the regionalized flood frequency models developed in the study. In some cases, however,

the RMSE is as high as 44% for regions having the steepest flood frequency growth curves indicating a lack of precision or fit. Paradoxically, this occurs with the GEV flood frequency model that should provide a better fit considering the fact that ot has one additional parameter to capture the high skew commonly found in flood data. By and large the RMSE is less than 20% for most cluster and USGS flood regions.

- 7. In both methods of delineating flood regions, the geomorphic properties of the watersheds such as the drainage area, basin shape, basin length, and main channel length, slope and sinuousity provide the maximum discrimination between flood regions. Discrimination is based on the physical and climatic characteristics of the watersheds (referred to as attribute variables). Watershed contributing drainage area is the most important variable. The USGS flood regions have a low overall discrimination (16.9%) compared to the cluster regions, which have a higher overall discrimination of 47.1%. This further supports the mixed hydrologic composition within the USGS regions.
- 8. For both methods of regionalization, the significant variables (at a 5% level) in the regression analysis relating EV1 and GEV flood quantiles, Q_{Ti} , of various return periods to watershed physical and climatic characteristics, are geomorphic properties of the watersheds (as was the case with discriminant analysis) with the watershed contributing drainage area, A_c , as the most important variable.
- 9. The standard errors associated with the regression equations are comparable for both methods. For cluster regions, where the problem of simultaneously classifying gauged and ungauged sites into several cluster regions exists, the weighted standard errors (based on the posterior probabilities and the

corresponding regression equations of the cluster regions to which a site is assigned) are used in making this comparison.

- 10. The hydrological characteristics of flood regions and their overall performance delineated using Method 1 (clustering on L-moments and QSP) are similar to those of Method 2 (clustering on parameters of the EV1 and GEV probability distributions). However, the actual gauged sites within each region are quite different.
- 11. A comparison of flood quantile estimates at selected sites indicates that the regionalized EV1 and GEV flood frequency models underestimate flood levels (50 and 100 year return periods) when compared to the log-Pearson Type-III flood frequency model.
- 12. Overall it appears that regionalized EV1 and GEV flood frequency models, in conjunction with the method of L-moments to estimate their parameters, would better represent flood experience in Kentucky even when using the present flood regions as defined using the method of residuals. This observation is based on the performance of these models in terms of bias and RMSE and not on a direct comparison with a regionalized log-Pearson Type-III flood frequency distribution.

NOMENCLATURE

The follo	wi	ng symbols and variables are used in this study:
LCV	=	L-moment ratio, t ₂ (coefficient of variation), of normalized AMF series;
lsk	=	L-moment ratio, t ₃ (coefficient of skewness), of normalized AMF series;
LKUR	=	L-moment ratio, t ₄ (coefficient of kurtosis), of normalized AMF series;
LBMD	Ξ	L-moment ratio, t ₅ (coefficient of bi-modality), of normalized AMF series;
CV	=	coefficient of variation of raw AMF series;
		coefficient of skewness of raw AMF series;
		coefficient of kurtosis of raw AMF series;
M(LCV)	=	regional weighted mean L-moment ratio, t ₂
		(coefficient of variation);
M(LSK)	=	regional weighted mean L-moment ratio, t _a
		(coefficient of skewness);
M(LKUR)	=	regional weighted mean L-moment ratio, t_A
		(coefficient of kurtosis);
M(LBMD)	=	regional weighted mean L-moment ratio, t ₅
		(coefficient of bi-modality);
AMF	=	raw or normalized annual maximum floodpeak
		series;
x	Ξ	normalized annual maximum flood value = Q/Q ;
Q	=	raw annual maximum flood value in cfs;
Q	=	mean of raw AMF series in cfs (same as STMEAN);
. Q _I	×	index-flood in cfs;
Q _{Ti}	=	gauged site i estimate of flood level having a
~ 		return period of T years in cfs.
$\mathtt{q}_{\mathbf{T}}$	=	regionalized flood quantile estimate of
-		normalized AMF series;

QSP = specific mean annual flood = Q/A_c in cfs/sq.

mile (same	as	SMDISCH);
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F = cumulative probability density function (cdf);

f = probability density function (pdf);

E(X) = expected value of random variable X;

 P_r^* (F) = rth shifted Legendre polynomial of function F;

- PWM = probability weighted moment;
- Mp,r,s = probability weighted moment;

 α_r = probability weighted moment of order r, $M_{1,0,r}$;

- β_r = probability weighted moment of order r, $M_{1,r,0}$;
- $\lambda_r = L$ -moment of order r;

 τ_r = L-moment ratio of order r;

 $a_r = sample estimate of PWM a_r;$

 b_r = sample estimate of PWM β_r ;

- l_r = sample estimate of the rth L-moment;
- $t_r = sample estimate of the rth L-moment ratio;$
- MEVL = sample estimate of the location parameter of the EV1 probability or frequency distribution;
- AEVL = sample estimate of the scale parameter of the EV1 probability or frequency distribution;
- MGVL = sample estimate of the location parameter of the GEV probability or frequency distribution;
- AGVL = sample estimate of the scale parameter of the GEV probability or frequency distribution;
- KGVL = sample estimate of the shape parameter of the GEV probability or frequency distribution;
 - EV1 = Extreme Value Type-1 or Gumbel probability
 distribution;

GEV = Generalized Extreme Value probability distribution;

WAK = Wakeby probability distribution;

 B_{tr} = watershed or basin width in miles;

B₁ = watershed or basin length in miles (same as BASLEN);

 B_{s} = watershed shape index = (A_{c} / B_{1}) (same as

SHAPE);

BELEV = watershed or basin mean elevation in feet;

PRECIP = watershed or basin mean annual precipitation in inches;

STOR = watershed or basin storage in percent;

SINFL = watershed or basin average soil infiltration in in/hr;

BASIN = watershed or basin designation;

S = main channel sinuousity = (L_c / B₁) (same as CHANSIN);

ELEV = gauged site mean elevation in feet;

ISTN = gauged site USGS Station Number;

T = return period in years;

N = number of years of AMF data at a gauged site;
USREG = region assigned to gauged site using method of residuals;

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APPENDIX A

TABLE A1. Important Statistics of Hydrological Characteristics Gauged Sites Used in the Study.

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0.613.0 0.2704 -0.0500 0.0091 1.245 74.191 0.4148 -</th><th>0.793 0.491 0.053 0.1400 2.471 3.840 0.410
0.411 0.347 -6.2503 0.3716 0.000 0.600 0.611</th><th>0.7532 0.4173 0.0245 0.1626 1.425 5.714 0.004 -</th><th></th><th>0.4344 0.4028 -6.1973 0.0000 21.703 70.007 0.7222 -
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0.4540 0.4023 -0.2110 0.1540 1.2730 0.2115 | 0.613.0 0.4144 -0.0500 0.1311 1.167 1.140 0.2173
0.613.0 0.2704 -0.0500 0.0091 1.245 74.191 0.4148 - | 0.793 0.491 0.053 0.1400 2.471 3.840 0.410
0.411 0.347 -6.2503 0.3716 0.000 0.600 0.611 | 0.7532 0.4173 0.0245 0.1626 1.425 5.714 0.004 - | | 0.4344 0.4028 -6.1973 0.0000 21.703 70.007 0.7222 -
0.4447 0.2221 -0.5440 0.4414 0.660 0.664 0.5714 | 1005-0 011-22 184"Y 0000-0 5051-0- 7211"B 6101-0 | 0767-9 681° - 4071 - 1927-9 547-9 140° - 140° - 2252° |
| SAUJSCH SHAFE MALER CHMIEN CHARSIN CHARSEN LCY | 0101.0 KT3.4 35M2.5 X1.1 K5.1 17.0 N2.3 K3.285
0 M11.0 1965.0 M11.0 X1.1 K5.1 N2.1 K1.1 K1.2 K2.1 K2.1 K1.1 K1.2 K2.1 K2.1
 | 1 11/1 11/1 11/1 11/1 11/1 11/1 11/1 1 | 1 4413-1 1141-1 1141-1 1121 1121 1141-1 1141
 | | | 1,005 2.170 11.00 2.200 1.401 0.101 0.112 0.123 0.001 0.002 0.1024 0.102 0.1021
 | 0 0441 9549 66542 2471 0623 672 672 672 677 1472
9 1673 11627 10967 5571 0054 079 0674 1473
0 1674 1627 16967 5571 0054 074 075 871 | AG'N MINY MIN' KIN MUL MUL MUL | 1000-0 102-0 500-0 100-0 510-0 100-0 100-0 100-0 102-0 100-00-00-00-00-00-00-00-00-00-00-00-00- | 6.4455 6.465 6.274 -0.2538 -0.2667 21.772 26.2742
6.2871 6.122 6.1251 0.1071 0.1041 1.426
6.2873 6.442 6.441 6.441 6.441 7.446 6.256 6.256 | MACO EX2 MACO E220- 1001-0- 2020 00210 2220 | 0.7232 0.4352 0.1426 -0.4735 -0.4521 3.444 E.555 0.3436 0.5528 0.5428 0.5528 0.5529 | - 1955, 0.111,0 0.5220 0.5000 0.6000 2.134 0.510 1.2597 -
12011 0.2000 0.3260 0.0255 0.5178 -0.2000 0.550 0.550
 | 9464-9421-64 220-42 0442-0-1042-0-1224-0-1224-0-2462-9-2462-0-242-0-2422-0-2422-0-2422-0-2422-0-2422-0-2422-0-2422-0-2422-0-2422-0-2422-0-242-0-2422-0-242-0-242-0-242-0-242-0-242-0-242-0-242-0-242-0-2422-0-242-0-242-0-242-0-242-0-242-0-242-0-242-0-242-0- | 0.5107 0.6134 0.444 -4.0508 0.1373 1.107 1.660 0.2173
0.6740 0.4154 6.2704 -6.6637 0.4098 13.746 74.674 0.4146 - | 0.3775 0.7775 0.7871 0.0750 0.1400 2.470 0.770 0.502
0.4573 0.4710 0.5517 0.2503 0.311 0.000 0.410 | 0.016 0.1532 0.073 0.8215 0.1421 1.425 5.714 0.006 - | 1200- 0.561 1.200 4.200 0.501 1.200 1.200 1.200 | 0.4730 0.4314 0.4028 -0.4873 0.0000 21.703 70.003 0.7222 -
6.6844 0.4442 0.2713 -0.5440 0.4414 0.664 0.664 0.5714 | 1005.6 011.55 101.0 0000.0 2020.0 1011.1 1100.0 501.0 | 076110 44112 40011 201210 5142-0 111710 5041-0 515210
225210 16514 81411 825110 854010- 515210 110810 515210
 |
| I BAREA SUBJECH SHAFE BABLER CHAMEN CHARSIN CHARGEP LEY | 0 0101.0 KTG.4 0500 Z.10 10.1 K1.0 K1.1 K1.0 K2.3 K1.0 K1.0 10.1 K1.0 K1.0 K1.0 K1.0 K1.0
 | 1 1954.00 34.011 [100] 11.30 54.200 2.401 [0.012] 54.201 64.201 [101] 101.00 10 | 1 14 14 14 14 14 14 14 14 14 14 14 14 14
 | | | 1 17-46 14 15 14 14 15 15 15 16 15 16 15 16 15 16 15 16 15 16 15 16 16 16 16 16 16 16 16 16 16 16 16 16
 | 0 | AG'N MINY MIN' KIN MUL MUL MUL | 0000 1727,0 2001 0.027 0.021 0.021 0.021 0.021 0.0200 0.0220 0.0200 0.0210 | 6.4455 6.465 6.274 -0.2538 -0.2667 21.772 26.2742
6.2871 6.122 6.1251 0.1071 0.1041 1.426
6.2873 6.442 6.441 6.441 6.441 7.446 6.256 6.256 | MACO EX2 MACO E220- 1001-0- 2020 00210 2220 | 0.7232 0.4352 0.1426 -0.4735 -0.4521 3.444 E.555 0.3436 0.5528 0.5428 0.5528 0.5529 | - 1955, 0.111,0 0.5220 0.5000 0.6000 2.134 0.510 1.2597 -
12011 0.2000 0.3260 0.0255 0.5178 -0.2000 0.550 0.550
 | 9464-9421-64 220-42 0442-0-1042-0-1224-0-1224-0-2462-9-2462-0-242-0-2422-0-2422-0-2422-0-2422-0-2422-0-2422-0-2422-0-2422-0-2422-0-2422-0-242-0-2422-0-242-0-242-0-242-0-242-0-242-0-242-0-242-0-242-0-2422-0-242-0-242-0-242-0-242-0-242-0-242-0-242-0-242-0- | 0.5107 0.6134 0.444 -4.0508 0.1373 1.107 1.660 0.2173
0.6740 0.4154 6.2704 -6.6637 0.4098 13.746 74.674 0.4146 - | 0.3775 0.7775 0.7871 0.0750 0.1400 2.470 0.770 0.502
0.4573 0.4710 0.5517 0.2503 0.311 0.000 0.410 | 0.016 0.1532 0.073 0.8215 0.1421 1.425 5.714 0.006 - | 1200- 0.554 1.204 -0.555 0.557 1.275 1.275 1.275 1.275 | 0.4730 0.4314 0.4028 -0.4873 0.0000 21.703 70.003 0.7222 -
6.6844 0.4442 0.2713 -0.5440 0.4414 0.664 0.664 0.5714 | 1005.6 011.55 101.0 0000.0 2020.0 1011.1 1100.0 501.0 | 076110 44112 40011 201210 5142-0 111710 5041-0 515210
225210 16514 81411 825110 854010- 515210 110810 515210
 |
| SAUJSCH SHAFE MALER CHMIEN CHARSIN CHARSEN LCY | 0101.0 KT3.4 35M2.5 X1.1 K5.1 17.0 N2.3 K3.285
0 M11.0 1965.0 M11.0 X1.1 K5.1 N2.1 K1.1 K1.2 K2.1 K2.1 K1.1 K1.2 K2.1 K2.1 | 1 1954.00 34.011 [100] 11.30 54.200 2.401 [0.012] 54.201 64.201 [101] 101.00 10 | 1 14 14 14 14 14 14 14 14 14 14 14 14 14 | | | 1 17-46 14 15 14 14 15 15 15 16 15 16 15 16 15 16 15 16 15 16 15 16 16 16 16 16 16 16 16 16 16 16 16 16 | 0 | AKM, AKM AKM AGM KGM, PART AUML MRT CINL | 9451'9 24C'9 24C'1 192'9 219'9 219'9 219'9 24C'1 192'9 24C'9 24C'1 192'9 | 12/12/2 (17/12) 12/12/2 (17/12) 12/12/2 (17/12) 12/12/2 (17/12) 12/12/2 (17/12) 12/12/2 (17/12) 12/12/2 (17/12) 12/12/2 (17/12) 12/12/2 (17/12) 12/12) 12/12/2 (17/12) 12/12/2 (17/12) 12/12) 12/12/2 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12 (17/12) 12/12) 12/12 (17/12) 12/12 | 1011 1111 1111 1111 1111 1111 1111 111 | 0.1242 0.2242 0.4722 0.1424 -0.4735 -0.4757 1.444 0.545 0.3434 0.4545 0.2542 0.4744 0.542 | - 1595, 9,452, 9,411,31 - 1522,9 - 1521,9 - 10,400 - 123,4 - 123,4 - 1597 - 1597,9 - | 0.045.0 0.51.0 0.52.0 0.52.0 0.50.0 0.52.0 0.52.0 0.51 | • 6.7651 0.5169 0.4736 0.4444 • 0.500 0.1571 1.169 0.5173
0.6563 0.6740 0.6174 0.5764 • 0.6037 0.6064 1.516 0.4146
- 0.6563 0.6740 0.6174 0.4746 | - 4.7421 9.3775 9.7978 9.4871 9.8758 9.1908 2.474 3.847 9.2402
6.7346 9.4573 8.4718 8.3413 -4.2503 8.3711 4.009 8.460 8.4173 | | CENT AND - 0.381 0.304 -0.302 0.302 0.403 1.403 1.403 | 6.7143 6.4756 6.4756 0.4028 -6.1872 0.0000 21.767 70.007 0.7222 -
• 6.3141 6.4644 6.4447 6.2713 -6.5446 0.4114 6.664 0.664 0.204 6.7114 | 1007 0 01122 101 0 000 0 5051 0- 1021 0 1100 0 1041 0 1045 0 | 0755-0 1817 - 001 - 1927-0 1949-0 1977-0 1947-0 1977-0 1977-0
2557-0 1817-1 - 1977-1 - 1977-0 1949-0 1977-0 1977-0 1977-0 1977-0 |
| II SINKER MARA SABISCH SHARE BARER CMMER CMARIN CHARGE LCY | 0 0101.0 KTG.4 0500 Z.10 10.1 K1.0 K1.1 K1.0 K2.3 K1.0 K1.0 10.1 K1.0 K1.0 K1.0 K1.0 K1.0
 | 1 1954.00 34.011 [100] 11.30 54.200 2.401 [0.012] 54.201 64.201 [101] 101.00 10 | 1 14 14 14 14 14 14 14 14 14 14 14 14 14
 | | | 1 17-46 14 15 14 14 15 15 15 16 15 16 15 16 15 16 15 16 15 16 15 16 16 16 16 16 16 16 16 16 16 16 16 16
 | 0 | KUM MEN MEN NEW FRA KAN KAN MAT KAN | | 2127 8 2218 22112 4907 9 2219 4219 1219 2219 2217 2217 2217 2217 2217 2 | 2011 2011 2011 2011 2011 1057 2011 2011 2011 2011 2011 2011 2011 201 | 0.234 0.427 | - 1442,1 610,4 6421,2 6444,0 6424,2 6424,3 6444,0 6424,9 6424,9 6424,1 - 1442,9 6424,9
6424,9 642 | 1.2112 0.782 0.784 0.721 0.2210 0.722- 1.727 0.2312 0.782 0.712 0.711 0.2112 0.781 0.2112 0.7810 0.7910 0.7910 | 0.1542 0.1814 0.1814 0.4814 0.4814 - 4.084 0.1914 1.181 1.181 0.4814 0.4814 - 4.414 - 4.414 0.481 | 2012'9 414'E 142'2 404'9 544'8 144'9 244'9 545'9 128'9 476'1
2012'9 409'9 909'9 312E'9 654'9 149'8 244'9 525'9 97E'9 551'9- | -4.505 0.7173 0.4169 0.7522 0.4173 0.0216 0.1626 1.425 1.425 5.714 0.4066 - | 2.2461 4.4664 4.3187 4.3261 4.3264 4.3254 4.4353 4.1578 4.453 4.453 | 6.7110 6.7113 6.4736 6.4746 0.4346 0.4628 -6.1817 0.0006 21.761 70.607 0.7222 -
2 5370 6.3191 6.4844 0.4447 6.2713 -1 5441 0.4414 0.669 0.660 0.3214 | 2007 0 0102 1001 0 000 0 1010 - 1111 0 1000 0 1000 1000 1000 | 01254 0121 0011 12121 014-5 1212 01210 12154 01215 01210 01210 01211 02221 02221 02221 02221 02221 02221 02221
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VSIGE IN STATEM BAREA SAUJSCH SAME BABLEN CHANTEN CHARSTAP LCY	0 0101.0 KTG3.0 350.055 GA1 10.0 KT 10	2 36 GATA 2010 MATAL LAN LLA SLAV 3.470 3.471 6.4245 6.227 6.141 6 2 61 DATA 145.00 24.241 1.52 31.41 5.106 2.124 6.4524 6.250 6.143 1 15 2412 1.542 24.2514 1.551 4.551 5.116 1.241 6.2224 6.457 6.2505 1 2 12 25.4 6.427 5.4253 1.241 1.551 1.241 1.552 7.226 6.229	No.1 No.1 <th< th=""><th></th><th></th><th>2 22 1043.47 177.46 16.463 2.176 17.16 23.26 1.161 1.161 1.161 1.171 1.171 1.271 1.1</th><th>0 - NALA 2014 2015 2017 002.2 1072 0072 1072 1072 10721 10721 10721 1072 9 11073 11173 20100 2017 0054 0195 0073 10722 00716 00551 11202 10 0 NALA 2017 10279 100179 2017 00742 0074 0175 0075 10751 10751 1120</th><th>KUM MEN MEN NEW FRA KAN KAN MAT KAN</th><th>9451'9 24C'9 24C'1 192'9 219'9 219'9 219'9 24C'1 192'9 24C'1 /th><th>2127 8 2218 22112 4907 9 2219 4219 1219 2219 2217 2217 2217 2217 2217 2</th><th>2011 2011 2011 2011 2011 1057 2011 2011 2011 2011 2011 2011 2011 201</th><th>0.234 0.427</th><th>- 1442,1 610,4 6421,2 6444,0 6424,2 6424,3 6444,0 6424,9 6424,9 6424,1 - 1442,9 6424,9 642</th><th>1.2112 0.782 0.784 0.721 0.2210 0.722- 1.727 0.2312 0.782 0.712 0.711 0.2112 0.781 0.2112 0.7810 0.7910 0.7910</th><th>0.1542 0.1814 0.1814 0.4814 0.4814 - 4.084 0.1914 1.181 1.181 0.4814 0.4814 - 4.414 - 4.414 0.481</th><th>2012'9 414'E 142'2 404'9 544'8 144'9 244'9 545'9 128'9 476'1 2012'9 409'9 909'9 312E'9 654'9 149'8 244'9 525'9 97E'9 551'9-</th><th>-4.505 0.7173 0.4169 0.7522 0.4173 0.0216 0.1626 1.425 1.425 5.714 0.4066 -</th><th>2.2461 4.4664 4.3187 4.3261 4.3264 4.3254 4.4353 4.1578 4.453 4.453</th><th>6.7110 6.7113 6.4736 6.4746 0.4346 0.4628 -6.1817 0.0006 21.761 70.607 0.7222 - 2 5370 6.3191 6.4844 0.4447 6.2713 -1 5441 0.4414 0.669 0.660 0.3214</th><th>2007 0 0102 1001 0 000 0 1010 - 1111 0 1000 0 1000 1000 1000</th><th>01254 0121 0011 12121 014-5 1212 01210 12154 01215 01210 01210 01211 02221 02221 02221 02221 02221 02221 02221</th></th<>			2 22 1043.47 177.46 16.463 2.176 17.16 23.26 1.161 1.161 1.161 1.171 1.171 1.271 1.1	0 - NALA 2014 2015 2017 002.2 1072 0072 1072 1072 10721 10721 10721 1072 9 11073 11173 20100 2017 0054 0195 0073 10722 00716 00551 11202 10 0 NALA 2017 10279 100179 2017 00742 0074 0175 0075 10751 10751 1120	KUM MEN MEN NEW FRA KAN KAN MAT KAN	9451'9 24C'9 24C'1 192'9 219'9 219'9 219'9 24C'1 192'9 24C'1	2127 8 2218 22112 4907 9 2219 4219 1219 2219 2217 2217 2217 2217 2217 2	2011 2011 2011 2011 2011 1057 2011 2011 2011 2011 2011 2011 2011 201	0.234 0.427	- 1442,1 610,4 6421,2 6444,0 6424,2 6424,3 6444,0 6424,9 6424,9 6424,1 - 1442,9 6424,9 642	1.2112 0.782 0.784 0.721 0.2210 0.722- 1.727 0.2312 0.782 0.712 0.711 0.2112 0.781 0.2112 0.7810 0.7910 0.7910	0.1542 0.1814 0.1814 0.4814 0.4814 - 4.084 0.1914 1.181 1.181 0.4814 0.4814 - 4.414 - 4.414 0.481	2012'9 414'E 142'2 404'9 544'8 144'9 244'9 545'9 128'9 476'1 2012'9 409'9 909'9 312E'9 654'9 149'8 244'9 525'9 97E'9 551'9-	-4.505 0.7173 0.4169 0.7522 0.4173 0.0216 0.1626 1.425 1.425 5.714 0.4066 -	2.2461 4.4664 4.3187 4.3261 4.3264 4.3254 4.4353 4.1578 4.453 4.453	6.7110 6.7113 6.4736 6.4746 0.4346 0.4628 -6.1817 0.0006 21.761 70.607 0.7222 - 2 5370 6.3191 6.4844 0.4447 6.2713 -1 5441 0.4414 0.669 0.660 0.3214	2007 0 0102 1001 0 000 0 1010 - 1111 0 1000 0 1000 1000 1000	01254 0121 0011 12121 014-5 1212 01210 12154 01215 01210 01210 01211 02221 02221 02221 02221 02221 02221 02221
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VSIGE IN STATEM BAREA SAUJSCH SAME BABLEN CHANTEN CHARSTAP LCY	0 0101.0 KTG.4 35MA45 STL1 K54.1 N.0 102.5 523 620,85 51.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0 1.0	1 101 0 1222 0 101 0 101 0 101 101 101 1	10012 1111 1111 1111 1111 1111 1111 1111 10012 1111 1111 1111 1111 1111 1111 1111 1111 10012 1111 1111 1111 1111 1111 1111 1111 1111 10012 1111 1111 1111 1111 1111 1111 1111 1111 10012 1111 1111 1112 1111 1112 1111 1111 1111 10013 1111 1111 1112 1111 1112 1111 1111 1111 10014 1111 1111 1112 1111 1111 1111 1111 1111 10014 1111 1111 1111 1111 1111 1111 1111 1111 10014 1111 1111 1111 1111 1111 1111 1111 1111 10014 1111 1111 1111 1111 1111 1111 1111 1111 10014 1111 1111 11111 11111 11111 11111 11111 11111 10014 1111 1111 1111 1111			CHUT 2 22 [MJJL] 177-06 14.065 2.179 15.66 2.200 1.261 0.1627 0.2112 9.1239 0 CHT 3 29 1.211.4 Re.06 11.211 1.2110 1.2109 1.2110 0.4012 0.4011 0.4023 CHT 3 10 1202.1 1.21 1.212 1.200 1.214 1.400 1.415 0.41012 0.2013 0.2014 0.4101 CHT 3 10 1202.1 1.01 11.125 1.230 1.214 1.400 1.425 1.41010 0.4101 0.2014 0.4012	0	ני נות אלא אלא אלא אלא אלא ונות גיאין אמנו אוגן אונין אונין אינין		2.42.04 7.3421 6.7427 6.4433 6.4642 6.274 - 4.257 - 6.2647 11.172 20.375 6.7342 6.7756 6.764 6.764 6.721 6.122 6.137 6.161 6.11 11.10 6.161 6.166 6.766 6.764 6.781 6.123 6.157 6.561 6.561 6.562 6.563	1,200 11,70 4,000 4,500 4,500 4,500 4,500 4,500 10,000 4,500	2,542/0 (4,212) 0.2423 0.2724 0.472 0.1244 0.125 0.4714 0.2513 0.2134 0.5424 0.2142 0.2529 0.4254 0.5424 0.5424 0.5424 0.5129 2.516 0.242	- 1485,1 CUC,4 MCL,5 MCM,9 MCL,8 SS21,8 MCL,9 MC		6.1278 6.1342 6.7183 9.5109 0.6143 0.4864 - 4.6500 0.1312 0.1573 5.1513 0.2143 0.2143 0.2143 0.2143 0.2143 0.2143 0.2144 0.2145 0.2144 0.2145	6.72121 1.41229 6.7121 9.7175 6.7192 6.4911 0.4954 9.1406 8.124 3.441 6.2422 1.4424 -6.4423 6.7346 6.4524 6.4510 6.3413 -6.2503 6.3711 0.006 6.400 6.40	LIGHT 4.2023 0.7339 0.1369 0.1532 0.4173 0.0266 0.2124 1.124 1.124 1.124 0.0266 -	2.2444 5.2441 4.4444 4.5341 4.5341 1.234 4.555 4.554 1.254 4.453 4.453	1.10753 0.2774 0.27143 0.4724 0.42344 0.4028 -0.1873 0.0006 21.767 70.007 0.2282 - 1.04451 2.4279 0.2141 0.4442 0.2243 -0.224 0.4444 0.644 0.644 0.644 0.444	2.5100 7.151 101.1 000.0 201.0. 121.1 0 101.0 10	0225-0 4872 4001 10222 0442-0 50270 10400 25570 2507.0 25671 102111 22527-0 48571 00511 02223 0440-0 5042-0 10400 6422-0 2517-0 25571 102111

TABLE AI (cont..). Important Statistics of Hydrological Characteristics of Gauged Sites Used in the Study. .

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TABLE A1 (cont..). Important Statistics of Hydrological Characteristics of Gauged Sites Used in the Study.

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STUISCA SAVE JASEN CHAREN CHARLEN CHARLEN LCY		. אפאר גופאר אאוגר קאונר איזור	
. MARA STATECH SWAE JASTEN CHMEN CHMENN CHWEND LCT	TAN 22.01 CAN CAN <thcan< t<="" th=""><th>. אפאר גופאר אאוגר קאונר איזור</th><th></th></thcan<>	. אפאר גופאר אאוגר קאונר איזור	
STUISCA SAVE JASEN CHAREN CHARLEN CHARLEN LCY		. אפאר גופאר אאוגר קאונר איזור	
JI ETIKEAR DAMEA BUDISCA SUWE DASLEN CHANELA CHANELAR LCY	TAN 22.01 CAN CAN <thcan< t<="" th=""><th>. NEW YEAR NOA VEAR LEAN WARE WARE MANE</th><th></th></thcan<>	. NEW YEAR NOA VEAR LEAN WARE WARE MANE	
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USING IN THICAN MANGA STANISCA SAWE BASLEN CHARLEN CHARSLOP LCY		. NEW YEAR NOA VEAR LEAN WARE WARE MANE	
MSIN BERGE IN ETHERN MARS STATISCA SAME ANSLET CAMEER CANASIN CHANSLOP LCY		אני נאש אניע אלא אניע אלא אינע אינער אינער אינער אינער אינער אינער אינער אינער אינער אינער אינער אינער אינער א	

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TABLE Al (cont..). Important Statistics of Hydrological Characteristics of Gauged Sites Used in the Study.

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TABLE A2.	Comparison (of	Regi	Lonal	Average	Historical	and Simulated
L-Moments:	Clustering (¢π	LCV	and	QSP. *		

Region No:			<u>Historic</u> <u>L-Mozents</u>		
3	1.0000	0.2383	0.1760	0.1810	0.0781
4	1.0000	0.2823	0.1988	0.1839	0.0786
1 2	1.0000	0.3242	0.2758	0.1914	0.1016
2	1.0000	0.3862	0.3058	0.1900	0.0809
5	1.0000	0.4432	0.4035	0.2784	0.1657
Region No:			age Simulate		
3	1.0000	0.2403	0.1781	0.1676	0.0653
4 .	1.0000	0.2816	0.1817	0.1644	0.0560
1	1.0000	0.3153	0.1910	0.1547	0.0717
2 · 5	1,0000	0.3536	0.2122	0.1488	0.0686
5	1.0000	0.3786	0.2239	0.1472	0.0789
Region			age Simulat		
<u>No:</u>		<u>L-Moz</u>	<u>ents using (</u>	GEN Z	
3	1.0000	0.2397	0.1870	0.1747	0.0722
4	1.0000	0.2796	0.2148	0.1883	0.0733
1	1.0000	0.3177	0.2739	0.2023	0.1047
2	1.0000	0.3633	0.2982	0.2033	0.1012
5	1.0000	0.4122	0.3655	0.2382	0.1348
	•				

 Regions arranged in increasing steepness of the corresponding flood frequency growth curves (i.e. increasing LCV or LSK)

TABLE A3. Comparison of Regional Average Historical and Simulated L-Moments: Clustering on EV1 Parameters and QSP. *

Region No:			<u>Mistoric</u> <u>L-Moments</u>		
4	1.0000	0.2310	0.1641	0.1829	0.0741
5 .	1,0000	0.2801	0.2008	0.1834	0.0743
	1.0000	0.3116	0.2637	0.1869	0.1001
2	1.0000	0.3621	0.2837	0.2074	0.1097
· 3 2 1	1.0000	0.4196	0.3703	0.2451	0.1359
<u>Region</u> No:			nce Simulat		
11974.					
4	1.0000	0.2329	0.1767	0.1687	0.0638
	1.0000	0.2798	0.1863	0.1694	0.0564
ā	1.0000	0.3053	0.1863	0.1554	0.0698
2	1.0000	0.3393	0.2078	0.1551	0.0638
5 3 2 1	1.0000	0.3678	0.2177	0.1477	0.0755
Region No:			nge Simulat		
4	1.0000	0.2328	0,1748	0.1730	0.0692
5	1.0000	0.2779	0.2175	0.1894	0.0716
3	1.0000	0.3057	0.2609	0.1976	0.1015
5 3 2 1	1.0000	0.3423	0.2673	0.1905	0.0870
1	1.0000	0.3950	0.3424	0.2305	0.1295

 Regions arranged in increasing steepness of the corresponding flood frequency growth curves (i.e. increasing LCV or LSK)

TABLE A4.	Comparison	of	Regional	Average	Historical	and	Simulated
L-Moments:	Clustering	on	LCV, LSKI	SW and QS	SP. *		

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Region No:			<u>Historic</u> <u>L-Moments</u>	• .	
4	1.0000	0.2229	0.0530	0.1324	0.0519
	1.0000	0.2613	0.2005	0.1713	0.0784
3 5 1 2	1,0000	0.3224	0.2420	0.1853	0.0809
ĩ	1.0000	0.3383	0.3187	0.2178	0.1176
2	1.0000	0.4615	0.4672	0.3275	0.1847
<u>Region</u> No:			and Simulat		
4	1.0000	0.2263	0.1785	0.1751	0.0615
	1.0000	0.2612	0.1791	0.1618	0.0679
J 5 1 2	1.0000	0.3109	0.1991	0.1604	0.0627
ĩ	1,0000	0.3240	0.1962	0.1545	0.0730
2	1.0000	0.3825	0.2274	0.1480	0.0817
<u>Region</u> No:			age Simulat ents using (
4	1.0000	0.2241	0.0886	0.1432	0.0391
3	1.0000	0.2604	. 0.2050 .	0.1748	0.0782
1 ·	1.0000	0.3296	0.3083	0.2198	0.1200
5	1.0000	0.3124	0.2465	0.1915	0.0847
2	1.0000	0.4221	0.4047	0.2669	0.1533

 Regions arranged in increasing steepness of the corresponding flood frequency growth curves (i.e. increasing LCV or LSX)

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Region No:			<u>Historic</u> <u>L-Moments</u>		
5	1.0000	0.2379	0.1054	0.1379	0.0527
4	1.0000	0.2735	0.0767	0.1239	0.0644
1	1.0000	0.2827	0.2877	0.2218	0.1145
1 2 3 5	1.0000	0.3330	0.3218	0.2319	0.0952
3	1.0000	0.3566	0.2335	0.1385	0.0700
6	1.0000	0.4233	0.4544	0.3236	0.1868
Region		Aver	age Simulat	ed	
Not			ents using		
5	1.0000	0.2399	0.1783	0.1685	0.0648
4	1.0000	0.2731	0.1889	0.1787	0.0543
1	1.0000	0.2802	0.1841	0.1593	0.0687
2	1.0000	0.3186	0.1961	0.1559	0.0660
3	1.0000	0.3363	0.2029	0.1535	0.0743
6	1.0000	0.3671	0.2185	0.1494	0.0766
Region		Aver	age Simulat	ed	
<u>Xo:</u>		L-Hom	ents using	GEV	
5	1.0000	0.2386	0.1284	0.1498	0.0516
4	1.0000	0.2701	0.1240	0.1476	0.0451
1	1.0000	0.2800	0.2804	0.2113	0.1086
3 2 6	1.0000	0.3419	0.2429	0.1782	0.0921
2	1.0000	0.3229	0.3022	0.2191	0.1034
6	1.0000	0.3983	0.4013	0.2682	0.1547

TABLE A5. Comparison of Regional Average Historical and Simulated L-Noments: Clustering on GZV Parameters and QSF. *

 Regions arranged in increasing steepness of the corresponding flood frequency growth curves (i.e. increasing LCV or LSK)

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Region Noi			<u>Historic</u> L-Moments		
6	1.0000	0.2698	0.2230	0.1867	0.0743
1	1.0000	0.2781	0.2728	0.2139	0.0989
4 2 7 3 5	1.0000	0.2830	0.2037	0.1582	0.0801
2	1.0000	0.2839	0.2265	0.1994	0.1044
7	1.0000	0.3034	0.2691	0.2028	0.1061
3	1.0000	0.3115	0.2695	0.1830	0.0887
5	1.0000	0.3185	0.2852	0.2061	0.1042
Region		Aver	age Simulat	ed.	
<u>Xo:</u>		L-Mom	ents using	5791	
6	1.0000	0.2622	0.1850	0.1648	0.0648
1	1.0000	0.2679	0.1887	0.1596	0.0669
4	1.0000	0.2770	0.1872	0.1626	0.0702
2 7 3 5	1.0000	0.2815	0,1881	0.1600	0.0682
7	1.0000	0.2959	0.1853	0.1564	0.0672
3	1.0000	0.3022	0.1909	0.1576	0.0689
5	1.0000	0.3040	0.1937	0.1577	0.0724
Region No:			age Simulat		
6	1.0000	0.2656			
7			0,2198	0.1904	0.0835
4	1.0000	0.2983	0.2666	0.2031	0.1026
	1.0000	0.2821	0.2182	0.1804	0.0865
2 1 3 5	1.0000	0.2844	0.2254	0.1845	0.0843
1	1.0000	0.2729	0.2597	0.2037	0.0986
3	1.0000	0.3028	0.2586	0.1934	0.0942
5	1.0000	0.3083	0.2727	0.2030	0.1041

TABLE A5. Comparison of Regional Average Historical and Simulated L-Moments: USGS Regions. *

Regions arranged in increasing steepness of the corresponding flood frequency growth curves (i.e. increasing LCV or LSK) ٠

TABLE A.7. Comparison of Regional Average Historic and Simulated Parameters: Clustering on LCV and QSP. *

Region	E.	71		GEV	
No. *	HEVI.	XEVL	Near	AGVL.	KGVL
		HISTORI	PARAMETERS		
3	0.80	0.34	0.40	0.34	-0.01
4	0.76	0.41	0.76	0.39	-0.04
1	0.73	0.47	0.70	0.40	-0.16
2 <u>.</u> 5	0.68	0.56	0.63	0.44	-0.20
5	0.63	0.64	0.55	0.42	-0.33
	l	VERAGE STHUL	TED PARAMETER	25	
3	0.80	0.35	0.80		
4	0.77	0.41	0.30	0.34	-0.03
1	0.74	0.45	0.71	0.38	-0.07
2	0.71	0.51	0.66		-0.16
S	0.68	0.55	0.59	0.42	-0.19
-		0.55	V.39	0.42	-0.25

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Regions arranged in increasing steepness of the corresponding flood frequency growth curves (i.e. increasing LCV or LSK) MEVL, MGVL = location parameters; AEVL, AGVL = scale parameters; and KGVL = shape parameter.

TABLE A8.	Comparison of Regional Average Historic and Simulated	
Parameters	: Clustering on EV1 Parameters and QSP. *	

		n		GEV		
Region No.*	HEVL.	XEVL	ngvi,	AGVL	KGVL	
		HISTORI	C PARAMETERS			
4	0.81	0.33	0.81	0.34	0.01	
5	0.77	0.40	0.76	0.39	-0.05	
5 3 2	0.74	0.45	0.71	0.39	-0.14	
2	0.70	0.52	0.66	0.43	-0.17	
1	0.65	0.61	0.58	0.43	-0.29	
	i	VERAGE SIMUL	ATED_PARAMETE	85		
4	0.81	0.34	0.81	0.33	-0.01	
5	0.77	0.40	0.76	0.37	-0.07	
5 3 2	0.75	0.44	0.72	0.38	-0.14	
2	0.72	0.49	0.69	0.42	-0.14	
1	0.69	0.53	0.62	0.42	-0.25	

Regions arranged in increasing steepness of the corresponding flood frequency growth curves (i.e. increasing LCV or LSX) MEVL, MGVL = location parameters; AEVL, AGVL = scale parameters; and KGVL = shape parameter.

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TABLE A9. Comparison of Regional Average Historics and Simulated 'Parameters: Clustering on LCV, LSKEN and QSP. *

B ==4==	E	71		GEV	
Region No. *	MEVL	XEVI.	MGVL	λgvi,	KGVL
		HISTORI	C PARAMETERS		
4	0.81	0.32	0.85	0.37	0.19
3 · 5	0.78	Q.38 T	0.77	0.36	-0.05
5	0.73	0.47	0.71	0.42	-0.11
1	0.72	0.49	0.68	0.38	-0.22
2.	0.62	0.67	9.52	0.38	-0.42
	· i	VERAGE STHUL	ATED PARAMETE	25	
4	0.81	0.33	0.83	0.36	0.13
3 5	0.78	0.38	0.77	0.36	-0.05
5	0.74	0.45	0.72	0.40	-0.12
1 .	0.73	0.47	0.69	0.38	-0.20
2	0.68	0.55	0.58	0.40	-0.34

Regions arranged in increasing steepness of the corresponding flood frequency growth curves (i.e. increasing LCV or LSK) MEVL, MGVL = location parameters; AEVL, AGVL = scale parameters; and KGVL = shape parameter.

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Baai an	:	271		GEV	
Region No.	HEVL.	AEVL	KGVL	ycar	XGVL
		HISTORIC	PARAMETERS		
5	0.80	0.34	0.82	0.37	0.10
4	0.77	0.39	0.80	0.45	0.15
1 2 3 6	0.76	0.41	0.74	0.34	-0.18
2	0.72	0.48	0.64	0.37	-0.22
3	0.70	0.51	0,68	0.47	-0.10
6	0.65	0.61	0.56	0.36	-0.40
	- 1	YERAGE STRUL	TTD PARAMETER	25	
s ,	0.80	0.35	0.81	0.36	0.07
4	0.77	0.39	0.79	0.41	0.07
1 2	0.77	0.40	0.74	0.34	-0.16
2	0.73	0.46	0.69	0.37	-0.20
3	0.72	0.49	0.69	0.44	-0.11
6	0.69	0.53	0.60	0.38	-0.33

TABLE AlG. Comparison of Regional Average Historic and Simulated Parameters: Clustering on G27 Parameters and QSP. «

Regions arranged in increasing steepness of the corresponding flood frequency growth curves (i.e. increasing LCV or LSX) MEVL, MGVL = location parameters; λ EVL, λ GVL = scale parameters; and KGVL = shape parameter. ٠ 1

Region No. +	EV1		GZY		
	NEAF	YEAL	MGVL	ygar	∦ KGVL
		HISTORI	C PARAMETERS		
6	0.78	0.39	9.76	0.36	-0.08
1	0.77	0.40	9.74	0.34	-0.15
4	0.76	0.41	0.75	0.39	-0.05
2	0.76	0.42	0.74	0.32	-0.09
7	0.75	0.44	0.72	0.37	-0.15
3	0.74	0.45	0.71	0.38	-9.15
5	9.73	0.46	0.70	0.38	-0.17
		AVERAGE STRUL	ATED_PARAMETER	<u>R5</u>	-
6	0.78	0.38	0.77	0.36	-0.08
1 .	0.78	0.39	0.75	0.34	-9.13
4	9.77	0.40	0.75	0.18	-0.07
2	9.77	0.41	0.75	0.38	-0.08
7	0.75	0.43	6.73	0.37	-0.14
3	0.75	0.44	0.72	0.38	-0.13
5	9.75	0.44	0.71	9.38	-0.15
				•	
				:	

TABLE All. Comparison of Regional Average Historic and Simulated Parameters: Clustering on USGS Regions. +

Frequency growth curves (i.e. increasing LCV or LSK) MEVL, MGVL = location parameters; AEVL, AGVL = scale parameters; and KGVL = shape parameter. ŧ

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