

The effects of diagenetic aromatization on the carbon and hydrogen isotopic composition of higher plant diand triterpenoids: evidence from buried wood

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1	The effects of diagenetic aromatization on the carbon and
2	hydrogen isotopic composition of higher plant di- and
3	triterpenoids: evidence from buried wood
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15	Abstract
16	The widely distributed aromatic di- and triterpenoids from higher plants occuring in
17	sediments are formed by diagenetic microbial processes affecting their precursor plant lipids,
18	and these compounds and their stable isotopic composition ($\delta^{13}C$ and $\delta^{2}H$) have the potential
19	to be used for palaeoenvironmental and palaeoclimatic studies. In the present study, the
20	isotopic composition of di- and triterpenoids aromatized to a different extent has been
21	measured to examine the isotopic effects associated with the aromatization reaction. To
22	overcome the possibility of multiple higher plant sources as is generally the case with
23	sedimentary lipids, the $\delta^{13}C$ and $\delta^{2}H$ values of aromatized di- and triterpenoids from lipid
24	extracts recovered from conifer and angiosperm buried wood have been determined, allowing
25	an unambiguous genetic precursor/product relationship to be insured. The results show that
26	the $\delta^{13}C$ compositions of both di- and triterpenes do not seem to be significantly affected by

progressive aromatization, whereas the situation is more contrasted with δ^2 H values. In the 27 case of diterpenoids related to abietic acid, a significant increase of the δ^2 H values by up to 86 28 29 ‰ with ongoing aromatization was measured. This is in contrast to what is expected for dehydrogenated compounds which should be globally more ²H-depleted than their precursor 30 molecules (e.g., biosynthesized unsaturated fatty acids vs. their saturated precursors). This ²H 31 32 enrichment of aromatized diterpenoids could indicate that they represent only minor residual 33 intermediates, the majority of these intermediates having been further degraded by processes favoring the degradation of the ¹H-containing substrates and having a moderately pronounced 34 $^{12}C/^{13}C$ selectivity. With triterpenoids, a preservation of the δ^2 H values was observed 35 whatever the nature and extent of the aromatization process considered, which may be related 36 37 to enzymatic reactions showing a limited carbon and hydrogen isotopic effect upon 38 aromatization. 39

Key words: Diterpenes, Triterpenes, Early diagenesis, Aromatization, Isotopes, Hydrogen,
Carbon, Buried wood.

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44 1. INTRODUCTION

The stable carbon (δ^{13} C) and hydrogen (δ^{2} H) isotopic composition of *n*-alkyl lipids from 45 plants (n-alkanes, n-alkanoic acids) in sediments have been actively investigated to acquire 46 47 plaeoenvironmental and palaeoclimatic information on, notably, the relative contributions of C3 and C4 plants (e.g., Huang et al., 2001; Liu et al., 2005; Tanner et al., 2007; Agrawal et al., 48 49 2014; Diefendorf and Freimuth, 2017; Rao et al., 2017) or on the reconstruction of past precipitation regimes or temperatures (e.g., Sauer et al., 2001; Sachse et al., 2004; Garcin et 50 51 al., 2012; Diefendorf and Freimuth, 2017; Freimuth et al., 2017; Rao et al., 2017). The same 52 palaeoenvironmental and palaeoclimatic information may potentially be provided by the determination of the δ^{13} C and δ^{2} H compositions of di- and triterpenic biomarkers from plants 53 54 in sediments. However, contrary to straight-chain lipids which occur in all plant species, di-55 and triterpenes can sometimes be associated to a specific plant familly or plant species (e.g., Jacob et al., 2008; Le Milbeau, 2010; Schnell et al., 2012, 2014). Investigation of their carbon 56 57 and hydrogen isotopic composition might therefore potentially provide additional and more 58 accurate information than that obtained from less or non-specific plant constituents. However, 59 terpenoids generally undergo alteration processes upon diagenesis, such as aromatization 60 which affects the structure of their genuine carbon skeleton and results in the loss of carbon and hydrogen atoms (Spyckerelle et al., 1977; Laflamme and Hites, 1979; Wakeham et al;, 61 62 1980; Wolff et al., 1989; Stout, 1992; Simoneit et al., 1986; Le Milbeau et al., 2010; Nakamura et al., 2010; Schnell et al., 2012, 2014). Since cleavage of several C-C and C-H 63 64 bonds might be associated to important carbon and hydrogen isotopic effects, a key question 65 arises whether the δ^{13} C and δ^{2} H values of aromatized terpenoids keep or not the original isotopic composition of their biological precursor molecules. This point is of particular 66 67 interest for studies in which the stable isotopic signatures of lipids are used. In addition, investigation of the carbon and hydrogen isotopic effects associated with aromatization might 68

also potentially give information about the mechanisms involved in these processes whichremain largely unknown.

71 To our knowledge, there are only two studies dealing specifically with the impact of early 72 diagenetic processes on the carbon (Freeman et al., 1994; Jacob et al., 2011) or hydrogen (Jacob et al., 2011) isotopic compositions of sedimentary aromatized triterpenoids from 73 angiosperms and one dedicated to the aromatisation processes affecting diterpenoids in coals 74 and mustones (Tuo et al., 2006). Nevertheless, comparison of the δ^{13} C or δ^{2} H values of di- or 75 76 triterpenoids aromatized to a different extent may occasionaly be found in articles not 77 specifically dealing with the impact of aromatization processes on carbon and hydrogen 78 isotopic composition, this aspect being generally not discussed (e.g., Schouten et al., 2007; 79 Tuo et al., 2003; Nakamura et al., 2010; Suzuki et al., 2010). The main problem while 80 investigating isotopic fractionation during aromatization of terpenes resides in the fact that a 81 given aromatized terpenoid within a sediment sample might originate from mixed plant sources. It results that the δ^{13} C and δ^{2} H differences between the aromatized intermediate 82 molecules within this sample might be due to fractionation effects associated with the 83 84 aromatisation process itself, but may also be related to source effects (i.e., contribution of 85 multiple source organisms with different isotopic compositions). To overcome this problem, 86 we have investigated the case of aromatized terpenoids within ancient buried woods samples. 87 They indeed represent ideal candidates to study the effects of early diagenetic aromatization 88 on carbon and hydrogen isotopic compositions of di- and triterpenoids since all the terpenoid 89 intermediates from a given buried wood sample must be unambiguously genetically related. 90 We report here the determination and possible significance of the variations of the δ^{13} C and 91 δ^2 H signatures of aromatic terpenoids from buried wood samples comprising conifer wood 92 (Pinus sylvestris) investigated for diterpenoids, and Quercus robur and Alnus sp. wood for 93 triterpenoids.

2. MATERIALS AND METHODS
2.1. Samples
Three buried wood samples were collected and analysed for their terpenoid content. They
comprise:
Sample 1: Piece of trunk of <i>Pinus sylvestris</i> buried under a mudslide (age : > 4000 BP; SE
France).
Sample 2 (cf. Schnell et al., 2014): Inner part of an oak trunk (heartwood) unearthed from a
palaeochannel of the Rhine river (¹⁴ C age: ca. 3700 BP; Gerstheim, NE France).
Sample 3 (cf. Schnell et al., 2014; Adam et al., 2014): Piece of alder heartwood (Alnus sp.)
from a palaeochannel of the Rhine river (¹⁴ C age: ca. 3700 BP; Gerstheim, NE France).
2.2. GC-MS
GC-MS analyses were performed on a Thermo Scientific Trace Ultra gas chromatograph
coupled to a Thermo Scientific TSQ Quantum mass spectrometer equipped with a Tri Plus
autosampler and a programmed temperature vaporizing (PTV) injector. MS conditions were:
source 220 °C, electron ionization (EI) at 70 eV and scanning m/z 50 to 700. GC conditions:
HP5-MS column (30 m x 0.25 mm, 0.1 μ m film thickness) using He as carrier gas (constant
flow, 1.1 ml/min); oven program: 70 °C-200 °C (10 °C/min), then to 300 °C (held 40 min) at
4 °C/min).

113 2.3. GC-irmMS

GC-irmMS measurements were carried out on a Delta V Plus mass spectrometer (Thermo
Scientific) coupled to a GC Trace GC Ultra gas chromatograph equipped with a Triplus

116	autosampler, an on-column injector, and an Agilent HP5-MS column (30 m x 0.25 mm i.d. x
117	$0.1 \ \mu m$ film thickness), and connected to a ConFlow IV interface system and a GC Isolink II
118	conversion unit, comprising a combustion oven at 1000 °C for δ^{13} C measurements (resp. a
119	pyrolysis oven at 1420°C for $\delta^2 H$ measurements). The temperature program was: 80 °C – 310
120	°C (4°C/min) – isothermal at 310 °C (40 min). Each analysis was repeated 3 x. Before and
121	after each triplicate, the carbon (resp. hydrogen) isotopic composition of a certified <i>n</i> -alkane
122	mixture (Type A5; Arndt Schimmelmann, Biogeochemical Laboratories, Indiana University,
123	USA) was measured and used for calibration. The stability of the measurements was checked
124	using pulses of reference CO ₂ (resp. H_2) prior (5 pulses) and after (3 pulses) each run. H_3^+
125	factor was determined every day. The data were analyzed using Isodat 3.0 software. For
126	correction of the δ^{13} C and δ^{2} H values of derivatized lipids, see section 2.7.
127	2.4. Extraction and fractionation of lipids from the pine wood sample
128	The buried pine tree wood sample was extracted with CH_2Cl_2/CH_3OH (1:1 v/v) according to
129	the procedure described by Schnell et al. (2014). 975 mg of total lipid extract (TLE) were
130	obtained from 5.52 g of wood, an aliquot of which (240 mg) being derivatized (1.
131	Ac ₂ O/Pyridine 1:1 v/v; 2. esterification with a solution of diazomethane in Et ₂ O, cf. Schnell et
132	al., 2014). The derivatized TLE was fractionated on a silica gel column yielding three
133	fractions eluted respectively with CH ₂ Cl ₂ /EtOAc (8:2 v/v) (F1, 168 mg; 2 x dead volume -D _v -
134); CH ₂ Cl ₂ /AcOEt (F2, 20.5 mg; 2 x D_v), and CH ₂ Cl ₂ /CH ₃ OH (1:1 v/v) (F3; 22.5 mg; 2 x D_v).
135	Fraction F1 was analysed by GC-MS.
136	Besides, the remaining underivatized extract (ca. 730 mg) was fractioned on a silica gel
137	column (D _v : 33 ml) to isolate diterpenoid hydrocarbon fractions for determination of their
138	carbon and hydrogen isotopic compositions. Eight fractions were recovered eluting,

139 respectively with cyclohexane (F'1 to F'5, 25 ml each), cyclohexane/CH₂Cl₂ (1:1 v/v, F'6 and

140 F'7, 25 and 30 ml respectively) and CH₂Cl₂/CH₃OH (1:1 v/v) (F'8; 50 ml). Fractions F'1-F'5

141 were analysed by GC-MS and GC-irmMS. Fraction F'8 was mainly composed of

142 tetrahydroabietic acid 1 and was further purified on a silica gel column using CH₂Cl₂/EtOAc

143 (95:5; v/v) as eluent, yielding almost pure tetrahydroabietic acid **1**. The latter was silylated

144 (section 2.7) prior analysis by GC-MS and GC-irmMS.

145 2.5. Fractionation of lipids from the oak wood sample

146 The sample from buried oak wood was extracted as described by Schnell et al. (2014). An

147 aliquot of the crude extract (40 mg) was fractionated on a silica gel column using CH₂Cl₂ as

eluent yielding, respectively, fraction F1 (11 mg; 3 D_v) and a more polar fraction F2 (5 mg; 3
D_v).

150 2.6. Fractionation of lipids from the alder wood sample

A moderately polar fraction (18 mg) obtained from a buried alder wood sample following a procedure involving extraction, derivatization and elution with CH₂Cl₂/EtOAc (8:2 v/v) on a silica gel column (cf. Adam et al.,2014) was further fractionated on a silica gel column using successively cyclohexane/CH₂Cl₂ (8:2; v/v) and EtOAc/CH₂Cl₂ (8:2; v/v) as eluents, yielding, respectively, an apolar fraction F1 (1.8 mg) containing the aromatic hydrocarbons which was further analyzed by GC-MS and GC-irmMS and a more polar fraction F2 (16 mg) not further investigated.

158 2.7. Derivatization

159 Fractions containing alcohols, phenols or carboxylic acids were derivatized prior GC-MS and

160 GC-irmMS analysis using pyridine (40 µl) and *bis-(N,O-trimethylsilyl)trifluoroacetamide*

161 (BSTFA; 150 μl, 70 °C, 2 h) having a known carbon and hydrogen isotopic composition of

162 the trimethylsilyl group (δ^{13} C : -41.14 ‰; δ^{2} H : -154.09 ‰). The solvent and excess reagent

163 were removed under a stream of N₂. δ^2 H and δ^{13} C values of derivatized biomarkers were 164 corrected according to Rieley (1994) for the added trimethylsilyl group.

165 **3**. RESULTS AND INTERPRETATIONS

166 3.1. Diagenetic aromatization of diterpenoids

167 GC-MS analysis of the organic extract of the sample of buried *Pinus sylvestris* (Figure 1) 168 revealed the presence of diagenetic transformation products of diterpenoid acids from the 169 abjetane and pimarane series such as tetrahydroabjetic acid 1, monounsaturated derivatives of 170 abietic and pimaric acids (2, 3) and dehydroabietic acid 4. The former was the predominant 171 constituent of the fraction and is considered as a typical microbial reduction product formed 172 under anaerobic conditions (Reunanen et al., 1990; Bailly et al., 2016). Beside diterpenic 173 acids, various saturated (fichtelite 5), monounsaturated (6), and aromatic hydrocarbons such 174 as 18-norabietatriene 7, tetrahydroretene 8 and retene 9 were detected. Such diterpenic 175 structures are frequently encountered in sediments with a significant contribution of biomass 176 originating from conifers (e.g., Wakeham et al., 1980; Simoneit et al., 1986; Hautevelle et al., 177 2006; Tuo et al., 2006; Nakamura et al., 2010; Suzuki et al., 2010). The aromatic compounds 178 7-9 are postulated to be the result of microbial aromatization processes operative under anoxic 179 conditions (Tavendale et al., 1997a,b; Martin et al., 1999).

For the investigation of the isotopic effects associated with aromatization processes, we have measured the carbon and hydrogen isotopic composition of the aromatic compounds **7**, **8**, and **9** after chromatographic separation of the total hydrocarbons from the TLE (see below). For comparison, the δ^2 H and δ^{13} C values of partly reduced compounds such as tetrahydroabietic acid **1**, one monounsaturated tricyclic hydrocarbon **6** and fichtelite **5** have been measured. This allowed us to evaluate specifically the influence of decarboxylation on the carbon isotopic composition of diterpenes (comparison between **1** and **5**). Because of partial

coelutions in GC, determination of the δ^2 H and δ^{13} C values of the various diterpenoids 187 188 discussed above could not be performed directly on the lipid fraction shown in Figure 1. 189 Therefore, further liquid chromatography fractionation was carried out, leading to several subfractions enriched in the diterpenoids of interest and allowing the δ^2 H and δ^{13} C compositions 190 191 of compounds 1 and 5-9 to be determined. 192 The results show that the carbon isotopic composition of the diterpenoids is retained, 193 independently of the diagenetic alteration process considered (decarboxylation, reduction, and 194 aromatization) (Figure 2). On the contrary, the δ^2 H signatures seem to be strongly affected by 195 these processes, which apparently result in a progressive increase of the δD values with 196 ongoing aromatization (from -238 ‰ for 18-norabietatriene 7 to -152 ‰ for retene 9). A 197 similar observation was made by Tuo et al. (2006) in the case of diterpenoids occurring in a 198 series of mudstone samples from the Liaohe Basin (China). Thus, the δ^2 H values determined 199 for dehydroabietane 10 by Tuo et al. (2006) were generally lower than those measured for the 200 related diaromatic diterpenoid simonellite 11 within the same samples, except in one case. 201 Hydrogenation processes affecting abietic acid and leading to the formation of 202 tetrahydroabietic acid 1, hydrocarbon 6 and fichtelite 5 in our wood sample (Figure 2) are 203 likely induced by microorganisms, which might explain the slightly lower δ^2 H values (ca. 25) % difference) between the hydrogenated compounds 1 and 5 and their unsaturated 204 205 counterpart (6). The D-depletion of the hydrogenated compounds can be explained by the fact 206 that the biological reactions involved in the reduction processes indeed favor the 207 incorporation of ²H-depleted hydrogen. Such a situation has been observed, for instance, by 208 Chikaraishi et al. (2009) for the biosynthesis of phytol from geranyl geraniol. By contrast, the 209 progressive increase of the δ^2 H values observed during aromatization of 18-norabietatriene 7

- (-238 ‰) to retene 9 (-152 ‰) is more difficult to interpret and will be discussed in detail in
 section 3.3.1.
- 212 3.2. Diagenetic aromatization of triterpenoids

213 3.2.1. Quercus robur

214 The triterpenoids recovered from the inner part of buried trunk of Q. robur unearthed from a 215 palaeochannel of the Rhine (NE, Gerstheim; France) have been investigated. A detailed 216 molecular analysis of the sample can be found in Schnell et al. (2014). The predominant 217 aromatized triterpenoids identified (12-17) bear an oxygen functionality at C-2 (Le Milbeau et 218 al., 2010; Schnell et al., 2012, 2014). These compounds have inherited this uncommon 219 structural feature from the predominant biological triterpenoid precursors 18 and 19 220 functionalized at both C-3 and C-2 which occur in fresh Q. robur wood. They result from 221 aromatization processes starting in ring D from the triterpenoid skeletons and progressing 222 towards ring A (Figure 3). Two chromatographic fractions, one enriched in the two 223 monoaromatic ketones 14-15 and tetraaromatic phenols 16-17 (Figure 4a) and the other 224 dominated by the alcohols 12 and 13 (Figure 4b) could be separated from the lipid extract of 225 the oak wood sample obtained according to the procedure described in Schnell et al. (2014). 226 These compounds correspond to two triplets of unambiguously genetically related 227 triterpenoids (Figure 3) which are ideal candidates for the evaluation of the effects of aromatization processes on both carbon and hydrogen isotopic compositions. The $\delta^{13}C$ and 228 229 δ^2 H values of triterpenoids 12-17 are given in Figure 3. Surprisingly, despite the loss of up to 230 six methyl groups from the precursor molecule 19 to its tetraaromatic phenol counterpart 17, 231 neither the carbon, nor the hydrogen isotopic compositions of the different intermediates seem 232 to be significantly affected by these aromatization processes progressing from ring D to ring 233 Α.

234 *3.2.2.* Alnus sp.

235 For comparison, the impact of aromatization processes on the carbon and hydrogen isotopic 236 compositions during aromatization of C-3-oxygenated triterpenes (as opposed C2,C3-237 difunctionalized precursor molecules) following the "classical" pathway starting from ring A 238 and progressing towards ring D (e.g., Wolff et al., 1989; Stout, 1992) has been investigated. 239 For that purpose, the carbon and hydrogen isotopic composition of aromatic triterpenoid 240 hydrocarbons related to lupanol 20 and β -amyrin 21 occurring in the organic extract of Alnus 241 *sp.* wood found at the same site as the oak wood sample have been determined. A detailed 242 analysis of the lipid extract of this sample is presented in Schnell et al. (2014) and Adam et al. 243 (2014). The partial gas chromatogram of the aromatic hydrocarbon fraction focusing on the 244 pentacyclic triterpenoids is presented in Figure 5. They comprise mono- (22) (Wolff et al., 1989), tri- (23, 24) and tetraaromatic (25) triterpenoids (Spyckerelle et al., 1977; Laflamme 245 246 and Hites, 1979; Wakeham et al., 1980; Chaffee & Fuchs, 1988), all being closely genetically related (Figure 6). Comparison of the δ^2 H and δ^{13} C values of the triterpenes from both 247 248 oleanane and lupane series showed that the hydrogen and carbon isotopic are not significantly 249 affected by the degree of aromatization, as was observed with the triterpenoids from the oak 250 sample (see above). Freeman et al. (1994) obtained similar results for carbon isotopes in the 251 case of aromatized triterpenoids occurring in Eocene oil shale samples (Messel formation, 252 Germany). In this case, however, the genetic relationship between the various aromatized 253 higher plant triterpenoids from the same series could not be guaranteed since the triterpenes likely originated from multiple higher plant sources. Regarding ${}^{2}H/{}^{1}H$ isotopes, there is only 254 255 one study which is devoted to the comparison of the δ^2 H composition of triterpenoids 256 aromatized to a different extent in sediments (Jacob et al., 2011). In this study, important 257 variations of up to 100 ‰ could be observed between potential triterpenoid precursors and 258 related aromatized compounds, a difference which we did not observe in our study. However,

as for the study of Freeman et al. (1994), the authors were dealing with triterpenoids extracted
from sediments, and for which multiple plant sources with different hydrogen isotopic
compositions can be expected.

262 3.3. Evolution of isotopic compositions upon aromatization

263 Early diagenetic aromatization processes observed in buried wood are triggered by anaerobic 264 microbial processes (e.g., Tavendale et al., 1997a,b; Harder and Foß, 1999; Martin et al., 265 1999) as demonstrated by incubation experiments involving either diterpenic acids 266 (Tavendale et al., 1997a,b) or triterpenoids (e.g., Trendel, 1985; Lohmann, 1988; Lohmann et 267 al., 1990). Therefore, since the formation of aromatized di- and triterpenoids is biologically-268 mediated, one could envisage that the carbon and hydrogen isotopic shifts induced by these 269 processes have some analogies with those from enzymatic reactions involved in the 270 dehydrogenation of fatty acids via desaturases (e.g., Behrouzian et al., 2001; Buist, 2004; 271 Buist and Behrouzian, 1998; Savile et al., 2001; Shanklin et al., 2009) or in the aromatization 272 of androgen steroids via aromatases (e.g., Simpson et al., 1994; Akhtar et al., 2011). Indeed, 273 similarly to the aromatization processes leading from sterols to aromatic steroid hormones 274 such as estradiol, the aromatization of diterpenoids and triterpenoids involves the loss of 275 angular methyl groups. Like for the aromatization of the ring A of androgen steroids, 276 terpenoid aromatization might proceed via the functionalization of the angular methyl groups. 277 Hydrogen isotopic effects related to the dehydrogenation reactions or loss of methyl groups 278 catalyzed by the enzymes involved have often been investigated in the frame of studies aimed 279 at elucidating their mechanisms, but the carbon isotope effects associated with these reactions 280 are generally not documented. Thus, for instance, in the case of fatty acid desaturases, the 281 hydrogen kinetic isotope effect has been shown to be strongly dependent on the class of 282 enzymes involved. This effect can be important in the case of membrane-bound desaturases 283 (e.g., Savile et al., 2001; Buist and Behrouzian, 1998), but is sometimes negligible with

284 soluble plant desaturases (e.g., Behrouzian et al., 2001; Buist, 2004). It should however be 285 noted that the enzymes involved in these biosynthetic pathways bear iron at their active site 286 and are O₂-dependent (Shanklin et al., 2009; Akhtar et al., 2011), being able to activate C-H 287 bonds aerobically. Therefore, since the early diagenetic aromatization of di- and triterpenes 288 apparently only occur under anaerobic conditions according to incubation experiments 289 (Trendel, 1985; Lohmann, 1988; Lohmann et al., 1990; Tavendale et al., 1997a,b; Harder and Foß, 1999; Martin et al., 1999), it is likely that the enzymatic processes initiating anaerobic 290 291 aromatization are significantly different and comprise C-H bond cleavage in the absence of 292 oxygen following different pathway(s). These pathways possibly involve the formation of a 293 radical by abstraction of an hydrogen atom as has been described for the enzymes responsible 294 for the biodegradation of hydrocarbons under strictly anaerobic conditions (e.g., Spormann 295 and Widdel, 2000; Buckel and Golding, 2006; Heider, 2007; Booker, 2009). Nevertheless, 296 regardless of the enzymatic pathway(s) involved (aerobic vs. anaerobic), the isotopic 297 selectivity in the initiation step of the aromatization processes can probably be expected to be similar in aerobic and anaerobic processes and to favor, globally, light isotopes (i.e., ¹H and 298 ${}^{12}C$ vs. ${}^{2}H$ and ${}^{13}C$). 299

300 3.3.1. Hydrogen isotopic composition

301 In the case of fatty acids biosynthesized by some marine macroalgae, it has been observed by 302 Chikaraishi et al. (2004) that the δ^2 H values of saturated fatty acids are significantly increased 303 compared to their unsaturated metabolites, which is probably related to the important 304 hydrogen isotopic effect associated to dehydrogenation reactions induced by membrane-305 bound desaturases (e.g., Savile et al., 2001; Buist and Behrouzian, 1998). In the case of 306 sequential dehydrogenation processes such as those leading from saturated fatty acids to 307 mono- di-, tri and tetra-unsaturated fatty acid analogues, the evolution of the δ^2 H values of the 308 intermediates was more difficult to predict. According to Chikaraishi et al. (2004), it appears

309 that these processes depend on the relative fluxes into the various unsaturated products and on 310 the conversion rates, since each dehydrogenation step affecting an intermediate is supposed to 311 result in a ²H-enrichment of this intermediate relative to the product formed. Similar 312 observations have been made when the hydrogen isotopic composition of alkenones was 313 compared to that of their related dehydrogenated products (D'Andrea et al., 2007; Schwab 314 and Sachs, 2009). It should, however, be noted that Zhang and Sachs (2007) observed a 315 different trend for the biosynthesis of C₁₈ unsaturated fatty acids from stearic acid by the 316 Chlorophyceae *Eudorina unicocca* and *Volvox aureus*. With these organisms, the δ^2 H values 317 measured for the C₁₈ unsaturated fatty acids were similar or slightly higher than those 318 determined for stearic acid. These authors explained this observation by the fact that the desaturases involved in the biosynthesis of unsaturated fatty acids in these organisms belong 319 320 to the so called "soluble" desaturases. As mentioned above, the action of these enzymes was 321 indeed shown to induce only very limited isotope effects (e.g., Behrouzian et al., 2001; Buist, 322 2004). In our case, the absence of a significant evolution of the δ^2 H value of the triterpenoids, 323 whatever the series considered (i.e., lupane vs. oleanane), the aromatization process (i.e., from 324 ring A to ring D or ring D to ring A; Figure 3 and Figure 6, respectively), and aromatization 325 progress (i.e., from mono- to tetraaromatic derivatives), might thus be explained by the fact 326 that the ${}^{1}H/{}^{2}H$ isotopic effect related to the bacterial enzymes involved in these aromatization 327 processes operating under oxygen-depleted conditions is limited, similarly to the "soluble" 328 desaturases reported from plants.

In the case of the aromatized diterpenoids from *Pinus silvestris*, the significant increase of the δ^2 H values observed with increasing aromatization is more difficult to explain in this context. One would indeed expect the aromatized diterpenes to have lower δ^2 H values or, as observed in the case of the triterpenoids, almost similar δ^2 H values than their precursor molecules. The present observation can only be explained if the aromatized diterpenoids identified

correspond to minor residual intermediates which have survived overall degradation processes 334 335 that led to further degradation products not detected in our study or to complete 336 remineralization. Such degradation processes would indeed most likely strongly favor the degradation of ¹H-enriched substrates, resulting in an increase of the δ^{2} H values of the 337 338 residual aromatic diterpenes. A similar δ^2 H trend has been reported in the case of the 339 anaerobic biodegradation of aromatic hydrocarbons such as toluene by denitrifying, sulfate-340 reducing and fermenting bacteria using the succinate pathway (e.g., Morash et al., 2001; 341 Meckenstock et al., 2004; Vogt et al., 2008). In this connection, it should be noted that the 342 δ^2 H values determined by Tuo et al. (2006) for dehydroabietane **10** in a series of mudstone samples from the Liaohe Basin were generally lower (except in one case) than those measured 343 for the related diaromatic diterpenoid simonellite **11** within the same samples. This overall 344 345 trend was thus close to that observed in our sample. Tuo et al. (2006) nevertheless proposed 346 another explanation for this trend, arguing that dehydrogenation processes most likely lead to 347 a ²H-enrichment of the products relative to the substrate, since hydrogen is preferentially 348 removed during such processes. This is, however, not the case as illustrated by the evolution 349 of the δ^2 H values of unsaturated lipids formed by dehydrogenation processes induced by 350 desaturases as compared to that of their saturated precursors (e.g., Chikaraishi et al., 2004; 351 D'Andrea et al., 2007; Schwab and Sachs, 2009). When a dehydrogenation reaction occurs, 352 the increase of the δ^2 H value of the remaining substrate relative to that of the unsaturated 353 product formed is due to the fact that the hydrogen isotopic effect mainly affects the hydrogen 354 atom at the C-H position involved in the in the first C-H cleavage reaction (e.g. Buist and 355 Behrouzian, 1996, 1998), this hydrogen atom involved in the desaturation being finally lost. 356 The δ^2 H value of the hydrogen atoms from the unsaturated product (or aromatized compound, in our case) formed should thus be almost identical to that of the substrate at the beginning of 357 358 the reaction, whereas that of the remaining unreacted substrate should progressively increase

359 (Figure 7). In the case of cascade reactions as observed for the diagenetic di- and 360 triterpenoids, it is of importance to note that each intermediate also corresponds to the 361 substrate of the next reaction, the same type of ²H isotope fractionation applying hence at each step. Furthermore, at each step, the extent of ²H isotopic fractionation will depend on the yield 362 363 of the reaction, the δ^2 H composition of the substrate increasing with increasing yield relative 364 to the product(s) formed. Therefore, given these considerations, it appears that the 365 interpretation of the ¹H/²H signatures of biomarkers formed by diagenetic processes are 366 extremely difficult to interpret in the case of palaeoenvironmental studies (e.g., past climatic 367 changes, palaeohydrology).

368 3.3.2. Stable carbon isotopic composition

369 The results show that the carbon isotopic compositions of both di- and triterpenes do not seem 370 to be significantly affected by progressive aromatization, regardless of the nature of the 371 substrate (di- vs. triterpenes), of the process involved (triterpene aromatization progressing 372 from ring A to ring D or from ring D to ring A) and of the number of carbon atom(s) lost 373 during the process (with up to 6 methyl groups lost in the case of tetraaromatic triterpenoids). 374 This suggests that the kinetics of the enzymatic processes for aromatization are not sensitive 375 to carbon isotopic effects. In the same way as that proposed to explain the evolution of the 376 δ^2 H signatures of terpenoids in relation with intense alteration processes (see above), one 377 might envisage that these degradation processes might potentially also lead to an increase of 378 the δ^{13} C values of the residual, undegraded aromatic terpenes. However, such a trend has not 379 been observed, and might be due to the fact that anaerobic biodegradation processes affecting 380 aromatic hydrocarbons can be significantly more sensitive to hydrogen isotopes than to 381 carbon isotopes, at least in some cases, as demonstrated for the anaerobic biodegradation of 382 toluene (e.g., Morash et al., 2001; Meckenstock et al., 2004; Vogt et al., 2008) or with 383 diterpenes (the present study).

384

385 4. CONCLUSIONS

386 The stable carbon and hydrogen isotopic composition of aromatic di- and triterpenoid 387 biomarkers from conifer and angiosperm buried wood has been investigated to determine the 388 extent of the isotopic fractionation involved in these diagenetic aromatization processes. Our 389 results indicate that the stable carbon isotopic composition is not significantly affected by aromatization regardless of the substrate (di- vs. triterpenoids), of the processes involved 390 391 (aromatization progressing from ring A to ring D or from ring D to ring A) and of the number 392 of carbon atom(s) lost during the process. This absence of changes in the terpenoid δ^{13} C 393 signatures may be attributed to the limited sensitivity of the enzymes involved in the 394 biological aromatization processes towards carbon isotopic effects, and indicates that $\delta^{13}C$ 395 signatures of aromatic higher plant lipids may thus be used with confidence to trace potential 396 biological sources (e.g., C3 vs. C4 plants).

397 In the case of the hydrogen isotopic composition of diagenetic terpenoids, our observations 398 are more contrasted, with the diterpenoid $\delta^2 H$ signatures showing a strong increase of the $\delta^2 H$ 399 values with ongoing aromatization, and triterpenoid $\delta^2 H$ composition remaining almost 400 constant whatever the nature and extent of the aromatization process involved. In the case of 401 diterpenoids, the trend observed could be attributed to the fact that the different aromatic 402 diterpenoids analyzed represent extremely minor intermediates which have survived to further 403 degradation/mineralization which affected the majority of the diterpenoid pool. In contrast, 404 the constancy of the triterpenoid $\delta^2 H$ signature suggests that ${}^{1}H/{}^{2}H$ isotopic effects related to 405 the bacterial enzymes involved in these aromatization processes operating under oxygen-406 depleted conditions is limited, similarly to what is reported for the "soluble" desaturases from plants. The high variability of the ¹H/²H values measured for aromatic terpenoids suggests 407

408	that the intern	oretation	of the	$^{1}H/^{2}H$	H signatures	of biomarker	s formed	by such	diagenetic

409 processes may be extremely difficult to interpret in connection with palaeoenvironmental

- 410 studies (e.g., climatic changes, palaeohydrology).
- 411

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418

419 REFERENCES

420 Adam, P., Schaeffer, P. Schmitt, G., Tardivon, H., Motsch, E., Schneider, N., Duringer, P.,

421 2014. Reinvestigation of the structure of a tetracyclic terpenoid from geological samples: A

422 novel marker for fungi involved in the degradation of buried wood. Organic Geochemistry 73,423 8-15.

424 Agrawal, S., Galy, V., Sanyal, P., Eglinton, T., 2014. C₄ plant expansion in the Ganga Plain

425 during the last glacial cycle: Insights from isotopic composition of vascular plant biomarkers.

- 426 Organic Geochemistry 67, 58-71.
- 427 Akhtar, M., Wright, J.N., Lee-Robichaud, P., 2011. A review of mechanistic studies on
- 428 aromatase (CYP19) and 17α-hydroxylase-17,20-lyase (CYP17). Journal of Steroid
- 429 Biochemistry and Molecular Biology 125, 2-12.

- 430 Bailly, L., Adam, P., Charrié, A., Connan, J., 2016. Identification of guaiacyl
- 431 dehydroabietates as novel markers of wood tar from Pinaceae in archaological samples.
- 432 Organic Geochemistry 100, 80-88.
- 433 Behrouzian, B., Buist, P.H., Shanklin, J., 2001. Application of KIE and thia approaches in the
- 434 mechanistic study of a plant steraoly-ACP Δ^9 desaturase. Chemical Communications 401-402.
- Booker, S.J., 2009. Anaerobic functionalization of unactivated C-H bonds. Current Opinion in
 Chemical Biology 13, 58-73.
- 437 Buckel, W., Golding, B.T., 2006. Radical enzymes in anaerobes. Annual Reviews in
- 438 Microbiology 60, 24-49.
- Buist, P.H., 2004. Fatty acid desaturase: selecting the dehydrogenation channel. Natural
 Product Reports 21, 249-262.
- Buist, P.H., Behrouzian, B., 1996. Use of deuterium kinetic isotope effects to probe
 the cryptoregiochemistry of Δ⁹ desaturation. Journal of the American Chemical Society 120,
 6295-6296.
- 444 Buist, P.H., Behrouzian, B., 1998. Deciphering the cryptoregiochemistry of oleate Δ^{12}
- 445 desaturase: a kinetic isotope effect study. Journal of the American Chemical Society 120,446 871-876.
- 447 Chaffee, A.L., Fuchs, C.J.R., 1988. Polycyclic aromatic hydrocarbons in Australian coals-III
- 448 Structural elucidation by proton magnetic resonance spectroscopy. Organic Geochemistry 12,449 261-271.

- 450 Chikaraishi, Y., Suzuki, Y., Naraoka, H., 2004. Hydrogen isotopic fractionations during
- 451 desaturation and elongation associated with polyunsaturated fatty acid biosynthesis in marine
- 452 macroalgae. Phytochemistry 65, 2293–2300.
- 453 Chikaraichi, Y., Tanaka, R., Tanaka, A., Ohkouchi, N., 2009. Fractionation of hydrogen
- 454 isotopes during phytol biosynthesis. Organic Geochemistry 40, 569-573.
- 455 D'Andrea, W.J., Liu, Z., Alexander, M.D.R., Wattey, S., Herbert, T.D., Huang, Y., 2007. An
- 456 efficient method for isolating individual long-chain alkenones for compound specific
- 457 hydrogen isotope analysis. Analytical Chemistry 79, 3430–3435.
- 458 Diefendorf, A.F., Freimuth, E.J., 2017. Extracting the most from terrestrial plant-derived n-
- 459 alkyl lipids and their carbon isotopes from the sedimentary record: A review. Organic
- 460 Geochemistry 103, 1-21.
- 461 Freeman, K.H., Boreham, C.J., Summons, R.E., Hayes, J.M., 1994. The effect of
- 462 aromatization on the isotopic compositions of hydrocarbons during early diagenesis. Organic
- 463 Geochemistry 21, 1037-1049.
- 464 Freimuth, E.J., Diefendorf, A.F., Lowell, T.W., 2017. Hydrogen isotopes of *n*-alkanes and *n*-
- 465 alkanoic acids as tracers of precipitation in a temperate forest and implications for
- 466 paleorecords. Geochimica et Cosmochimica Acta 206, 166-183.
- 467 Garcin, Y., Schwab, V.F., Gleixner, G., Kahmen, A., Todou, G., Séné, O., Onana, J.M.,
- 468 Achoundong, G., Sachse, D., 2012. Hydrogen isotope ratios of lacustrine sedimentary n-
- 469 alkanes as proxies of tropical African hydrology: Insights from a calibration transect across
- 470 Cameroon. Geochimica et Cosmochimica Acta 79, 106-126.

- 471 Harder, J., Foss, S., 1999. Anaerobic formation of aromatic hydrocarbon p-cymene from
- 472 monoterpenes by methanogenic enrichment cultures. Geomicrobiology Journal 16, 295-305.
- 473 Hautevelle, Y., Michels, R., Lannuzel, F., Malartre, F., Trouiller, A., 2006. Vascular plant
- 474 biomarkers as proxies for palaeoflora and palaeoclimatic changes at the Dogger/Malm
- transition of the Paris Basin (France). Organic Geochemistry 37, 610–625.
- 476 Heider, J., 2007. Adding handles to unhandy substrates : anaerobic hydrocarbon activation
- 477 mechanisms. Current Opinion in Chemical Biology 11, 188-194.
- 478 Huang, Y., Street-Perrott, F.A., Metcalf, S.E., Brenner, M., Moreland, M., & Freeman, K.H.
- 479 (2001). Climate change as the dominant control on glacial-interglacial variations in C_3 and C_4
- 480 plant abundance. *Science*, 293, 1647-1651.
- 481 Jacob, J., Disnar, J.R., Arnaud, F., Chapron, E., Debret, M., Lallier-Vergès, E., Desmet, M.,
- 482 Revel-Rolland, M., 2008. Millet cultivation history in the French Alps as evidenced by a
- 483 sedimentary molecule. Journal of Archaeological Science 35, 814-820.
- 484 Jacob, J. Le Milbeau, C., Bossart, N., Disnar, J.R., & Billaud, Y., 2011. Combined δ^{13} C δ D
- 485 analysis of pentacyclic triterpenes and their derivatives, In: Dieckman, V., Tegelaar, R., van
- 486 Bergen, P., Almendros, G., Bernasconi, S., Schmidt, M., Schubert, C. (Eds.), The 25th
- 487 International Meeting on Organic Geochemistry. Book of Abstracts. Rapiergroup, Hampton
 488 Hill, p. 273.
- 489 Laflamme, R.E., Hites, R.A., 1979. Tetra- and pentacyclic, naturally-occurring, aromatic
- 490 hydrocarbons in recent sediments. Geochimica et Cosmochimica Acta 43, 1687-1692.

- 491 Le Milbeau, C., Schaeffer, P., Connan, J., Albrecht, P., Adam, P., 2010. Aromatized C-2
- 492 oxygenated triterpenoids as indicators for a new transformation pathway in the environment.
- 493 Organic Letters 12, 1504-1507.
- 494 Liu, W., Huang, Y., An, Z., Clemens, S.C., Li, L., Prell, W.L., Ning, Y., 2005. Summer
- 495 monsoon intensity controls C₄/C₃ plant abundance during the last 35 ka in the Chinese Loess
- 496 Plateau : Carbon isotope evidence from bulk organic matter and individual leaf waxes.
- 497 Palaeogeography, Palaeoclimatology, Palaeoecology 220, 243-254.
- 498 Lohman, F., 1988. Aromatisations microbiennes de triterpènes végétaux. PhD thesis,
- 499 Université Louis Pasteur, Strasbourg, France.
- 500 Lohmann, F., Trendel, J.M., Hetru, C., Albrecht, P., 1990. C-29 tritiated β-amyrin : chemical
- 501 synthesis aiming at the study of aromatization processes in sediments. Journal of Labelled
- 502 Compounds and Radiopharmaceuticals 28, 377-386.
- 503 Martin, V.J., Yu, Z., Mohn, W.W., 1999. Recent advances in understanding resin acid
- 504 biodegradation: microbial diversity and metabolism. Archives of Microbiology 172, 131-138.
- 505 Meckenstock, R.U., Morasch, B., Griebler, C., Richnow, H.H., 2004. Stable isotope
- 506 fractionation analysis as a tool to monitor biodegradation in contaminated aquifers. Journal of
- 507 Contaminant Hydrology 75, 215-255.
- 508 Morasch, B., Richnow, H.H., Schink, B., Meckenstock, R.U., 2001. Carbon and hydrogen
- 509 stable isotope fractionation during microbial toluene degradation: Mechanistic and
- 510 environmental aspects. Applied and Environmental Microbiology 67, 4842-4849.

- 511 Nakamura, H., Sawada, K., & Takahashi, M., 2010. Aliphatic and aromatic terpenoid
- 512 biomarkers in Cretaceous and Paleogene angiosperm fossils from Japan. Organic
- 513 Geochemistry 41, 975-980.
- 514 Rao Z., Guo, W., Cao, J., Shi, F., Jiang, H., & Li, C., 2017. Relationship between the stable
- 515 isotopic composition of modern plants and surface soils and climate: A global review. Earth-
- 516 Science Reviews 165, 110-119.
- Reunanen, M., Ekman, R., & Heinonen, M., 1990. Long term alteration of pine tar in a marine
 environment. Holzforschung 44, 277–278.
- 519 Rieley, G., 1994. Derivatization of organic compounds prior to gas chromatographic-
- 520 combustion-isotope ratio mass spectrometric analysis: identification of isotope fractionation
- 521 processes. Analyst 119, 915-919.
- 522 Sachse, D., Radke, J., Gleixner, G., 2004. Hydrogen isotope ratios of recent lacustrine
- sedimentary *n*-alkanes record moderne climate variability. Geochimica et Cosmochimica Acta
 68, 4877-4889.
- 525 Savile, C.K., Reed, D.W., Meesapyodsuk, D., Covello, P.S., Buist, P.H., 2001.
- 526 Cryptochemistry of a *Brassica napus* fatty acid desaturase: a kinetic isotope effect. Journal of
- 527 the Chemical Society, Perkin Transactions I 1116-1121.
- 528 Sauer, P., Eglinton, T.I., Hayes, J.M., Schimmelmann, A., Sessions, A., 2001. Compound-
- 529 specific D/H ratios of lipid biomarkers from sediments as a proxy for environmental and
- 530 climatic conditions. Geochimica et Cosmochimica Acta 65, 213-222.

- 531 Schnell, G., Schaeffer, P., Motsch, E., Adam, P., 2012. Triterpenoids functionalized at C-2 as
- 532 diagenetic transformation products of 2,3-dioxygenated triterpenoids from higher plants in
- 533 buried wood. Organic & Biomolecular Chemistry 10, 8276-8282.
- 534 Schnell, G., Schaeffer, P., Tardivon, H., Motsch, E., Connan, J., Ertlen, D., Schwartz, D.,
- 535 Schneider, N., Adam, P., 2014. Contrasting diagenetic pathways of higher plant triterpenoids
- in buried wood as a function of tree species. Organic Geochemistry 66, 107-124.
- 537 Schouten, S., Woltering, M., Rijpstra, W.I.C., Sluijs, A., Brinkhuis, H., Sinninghe-Damsté,
- 538 J.S.S., 2007. The Paleocene-Eocene carbon isotope excursion in higher plant organic matter:
- 539 Differential fractionation of angiosperms and conifers in the Arctic. Earth and Planetary
- 540 Science Letters 258, 581-592.
- 541 Schwab, V.F., Sachs, J.P., 2009. The measurement of D/H ratio in alkenones and their
- 542 isotopic heterogeneity. Organic Geochemistry 40, 111–118.
- Shanklin, J., Guy, J.E., Mishra, G., & Lindquist, Y., 2009. Desaturases : Emerging models for
 understanding functional diversification of diiron-containing enzymes. Journal of Biological
 Chemistry 284, 18559-18563.
- 546 Simoneit, B.R.T., Grimalt, J.O., Wang, T.G., Cox, R.E., Hatcher, P.G., Nissenbaum, A.,
- 547 1986. Cyclic terpenoids of contemporary resinous plant detritus and of fossil woods, ambers
- and coals. Organic Geochemistry 10, 877–889.
- 549 Simpson, E.R., Mahendro, M.S., Means, G.D., Kilgore, M.W., Hinshelwood, M.M., Graham-
- 550 Lorence, S., Amarneh, B., Ito, Y., Fisher, C.R., Dodson Michael, M., Mendelson, C.R.,
- 551 Bulun, S.E., 1994. Aromatase cytochrome P450, the enzyme responsible for estrogen
- 552 biosynthesis. Endocrine Reviews 15, 342-355.

- 553 Spyckerelle, C., Greiner, A., Albrecht, P., Ourisson, G., 1977. Aromatic hydrocarbons from
- 554 geological sources III. A tetrahydrochrysene derived from triterpenes in recent and old
- sediments: 3,3,7-trimethyl-1,2,3,4-tetrahydrochrysene. Journal of Chemical Research (M)

556 3746-3777.

- 557 Spormann, A.M., Widdel, F., 2000. Metabolism of alkylbenzenes, alkanes, and other
- 558 hydrocarbons in anaerobic bacteria. Biodegradation 11, 85-105.
- 559 Stout, S.A., 1992. Aliphatic and aromatic triterpenoid hydrocarbons in a Tertiary
- angiospermous lignite. Organic Geochemistry 18, 51-66.
- 561 Suzuki, N., Yessalina, S., Kikuchi, T., 2010. Probable fungal origin of perylene in Late
- 562 Cretaceous to Paleogene terrestrial sedimentary rocks of northeast Japan as indicated from

stable carbon isotopes. Organic Geochemistry 41, 234-241.

- 564 Tanner, B.R., Uhle, M.E., Kelley, J.T., Mora, C.I., 2007. C3/C4 variations in salt-marsh
- 565 sediments: An application of compound specific isotopic analysis of lipid biomarkers to late

566 Holocene paleoenvironmental research. Organic Geochemistry 38, 474-484.

- 567 Tavendale, M.H., McFarlane, P.N., Mackie, K.L., Wilkins, A.L., Langdon, A.G., 1997a. The
- 568 fate of resin acids-1. The biotransformation and degradation of deuterium labelled

569 dehydroabietic acid in anaerobic sediments. Chemosphere 35, 2137-2151.

- 570 Tavendale, M.H., McFarlane, P.N., Mackie, K.L., Wilkins, A.L., Langdon, A.G. 1997b. The
- 571 fate of resin acids-2. The fate of resin acids and resin acid derived neutral compounds in
- anaerobic sediments. Chemosphere 35, 2153-2166.
- 573 Trendel, J., 1985. Dégradation de triterpènes dans les sédiments. Aspects photochimiques et
- 574 microbiologiques. PhD thesis, Université Louis Pasteur, Strasbourg, France.

- 575 Tuo, UJ., Wang, X., Chen, J., & Simoneit, B.R.T., 2003. Aliphatic and diterpenoid
- 576 hydrocarbons and their individual carbon isotopic composition in coals from the Liaohe
- 577 Basin, China. Organic Geochemistry 34, 1615-1625.
- 578 Tuo, J., Zhang, M., Wang, X., Zhang, C., 2006. Hydrogen isotope ratios of aliphatic and
- 579 diterpenoid hydrocarbons in coals and carbonaceous mudstones from the Liaohe Basin,
- 580 China. Organic Geochemistry 37, 165-176.
- 581 Vogt, C., Cyrus, E., Herklotz, I., Schlosser, D., Bahr, A., Herrmann, S., Richnow, H.H.,
- 582 Fischer, A., 2008. Evaluation of toluene degradation pathways by two-dimensional stable
- 583 isotope fractionation. Environmental Science & Technology 42, 7793-7800.
- 584 Wakeham, S.G., Schaffner, C., Giger, W., 1980. Polycyclic aromatic hydrocarbons in Recent
- 585 lake sediments-II Compounds derived from biogenic precursors during early diagenesis.
- 586 Geochimica et Cosmochimica Acta 44, 415-430.
- 587 Wolff, G., Trendel, J.M., Albrecht, P., 1989. Novel monoaromatic triterpenoid hydrocarbons
- 588 occurring in sediments. Tetrahedron 45, 6721-6728.
- Zhang, Z., Sachs, J.P., 2007. Hydrogen isotope fractionation in freshwater algae: I Variation
 among lipids and species. Organic Geochemistry 38, 582-608.
- 591

592 FIGURE CAPTIONS

FIGURE 1 Partial gas chromatogram (GC-MS) of the apolar fraction F1 (cf. section 2.4) from
the TLE recovered from buried pine wood (SE France). Carboxylic acids were analysed as
methyl esters.

596 FIGURE 2 Scheme showing the genetic relationship between abietic acid and its related

597 diagenetic diterpenoids occurring in buried pine wood. Measured δ^2 H and δ^{13} C values are

598 indicated under the corresponding structure. In the case of tetrahydroabietic acid 1, the

599 isotopic values were measured on the related trimethylsilyl derivative and were corrected for

600 the added derivatization group introduced (cf. section 2.7). **a**: aromatization; **b**:

601 hydrogenation; **c**: decarboxylation.

602 FIGURE 3 Scheme showing the genetic relationship between aromatic C-2 oxygenated

triterpenoids 12-17 occurring in buried oak wood and bartogenic 18 and 23-

604 hydroxybartogenic 19 acids (adapted from Schnell et al., 2012 and 2014). Measured δD and

 δ^{13} C values are indicated under the corresponding structure. Alcohols **12-13** and phenols **16**-

606 17 were analysed as trimethylsilyl derivatives and the δ^2 H and δ^{13} C values were corrected for

607 the derivatization group introduced (cf. § 2.7).

608 FIGURE 4 Partial gas chromatograms (GC-MS) of fraction (a) F1; (b) F2 (cf. section 2.5)

609 isolated from the TLE from a buried oak (Q. robur) trunk (heartwood) sample collected in a

610 palaeochannel from the Rhine river (Gerstheim, NE France). Alcohols and phenols were

611 analysed as trimethylsilyl derivatives.

612 FIGURE 5 Partial gas chromatograms (GC-MS) of the apolar fraction F1 (cf. section 2.6)

613 isolated from the lipid extract from a buried alder heartwood (Alnus sp.) sample collected in a

614 palaeochannel from the Rhine river (Gerstheim, NE France).

615 FIGURE 6 Scheme showing the genetic relationship between aromatic triterpenoids 22-25

616 occurring in buried alder wood found in Gerstheim (NE France). Measured δ^2 H and δ^{13} C

617 values are indicated under the corresponding structures.

618 FIGURE 7 Scheme summarizing the evolution of the hydrogen isotopic composition of the

619 monounsaturated transformation product and the residual substrate during early diagenetic

620 dehydrogenations. Biologically triggered dehydrogenation reactions are generally sensitive to 621 the isotopic composition of the hydrogen atom (H_{1a biol}) involved in the first C-H cleavage 622 reaction and not to that involving the second hydrogen atom (H_{2a biol}) removed during the 623 dehydrogenation step (e.g., Buist and Behrouzian, 1996; 1998). The role on the isotopic 624 composition of the hydrogen atoms not involved in the reaction should be negligible 625 (secondary isotopic effects). It results that the δ^2 H value of the remaining substrate should increase mainly due to the progressive ²H-enrichement of the hydrogen at the position 626 627 involved in the first C-H cleavage reaction whereas the δ^2 H value of the monounsaturated 628 transformation product should be close to that of the original biological substrate prior 629 dehydrogenation (unless the latter serves as a substrate for further dehydrogenation). 630

Appendix





Fig. 1



Fig. 2



Fig. 3



Fig. 4



Fig. 5







Fig. 7