

Mapping blue-ice areas in Antarctica using ETM+ and MODIS data

Fengming HUI,¹ Tianyu CI,¹ Xiao CHENG,¹ Ted A. SCAMBOS,² Yan LIU,¹
Yanmei ZHANG,³ Zhaohui CHI,⁴ Huabing HUANG,¹ Xianwei WANG,¹
Fang WANG,¹ Chen ZHAO,¹ Zhenyu JIN,^{1,5} Kun WANG¹

¹State Key Laboratory of Remote Sensing Science, and College of Global Change and Earth System Science, Beijing Normal University, Beijing, China

E-mail: xcheng@bnu.edu.cn

²National Snow and Ice Data Center, CIRES, University of Colorado, Boulder, CO, USA

³Institute of Earthquake Science, China Earthquake Administration, Beijing, China

⁴Department of Geography, Texas A&M University, TX, USA

⁵Department of Geography, University of Utah, Salt Lake City, UT, USA

ABSTRACT. Blue-ice areas (BIAs) and their geographical distribution in Antarctica were mapped using Landsat-7 ETM+ images with 15 m spatial resolution obtained during the 1999–2003 austral summers and covering the area north of 82.5°S, and a snow grain-size image of the MODIS-based Mosaic of Antarctica (MOA) dataset with 125 m grid spacing acquired during the 2003/04 austral summer from 82.5°S to the South Pole. A map of BIAs was created with algorithms of thresholds based on band ratio and reflectance for ETM+ data and thresholds based on snow grain size for the MOA dataset. The underlying principle is that blue ice can be separated from snow or rock by their spectral discrepancies and by different grain sizes of snow and ice. We estimate the total area of BIAs in Antarctica during the data acquisition period is 234 549 km², or 1.67% of the area of the continent. Blue ice is scattered widely over the continent but is generally located in coastal or mountainous regions. The BIA dataset presented in this study is the first map covering the entire Antarctic continent sourced solely from ETM+ and MODIS data. This dataset can potentially benefit other studies in glaciology, meteorology, climatology and paleoclimate, meteorite collection and airstrip site selection.

KEYWORDS: Antarctic glaciology, blue ice, remote sensing

INTRODUCTION

Blue-ice areas (BIAs) account for ~1% of the Antarctic surface area (Bintanja, 1999; Winther and others, 2001) and are widely scattered over the continent, though they are generally located in coastal or mountainous regions. In spite of their small relative area, BIAs are important in many studies because of their surface characteristics: they are flat, smooth and hard. Very old ice is exposed at the surface of wind-induced BIAs, which occur near mountains or on outlet glaciers at higher elevations. In these areas, persistent winds erode snow year-round and combine with sublimation to create areas of net ablation (Winther and others, 2001). Studies have indicated that sampling ice in these areas could provide higher temporal resolution than is found in deep ice cores (Moore and others, 2006; Sinisalo and others, 2007). Consequently, BIAs are of great potential interest for paleoclimate studies (Bintanja, 1999; Grinsted and others, 2003; Moore and others, 2006; Ackert and others, 2011; Korotkikh and others, 2011; Fogwill and others, 2012; Spaulding and others, 2012). A detailed review of Antarctic BIAs with emphasis on their paleoclimate applications has been conducted (Sinisalo and Moore, 2010). The albedo of BIAs in different places in Antarctica has been summarized (Warren and Brandt, 2006) and found to range from 0.55 to 0.66 over the continent (Weller, 1980; Bintanja and Van den Broeke, 1995; Bintanja and others, 1997; King and Turner, 1997; Liston and others, 1999; Reijmer and others, 2001; Warren and others, 2002). The lower albedo of BIAs could influence the local surface energy balance and climate (Bintanja and Van den Broeke,

1994, 1995; Van den Broeke and Bintanja, 1995; Bintanja and others, 1997; Bintanja, 2000; Rasmus and Beckmann, 2007). BIAs generally occur in ablation areas, where there is a negative surface mass balance because annual ablation exceeds accumulation, removing the snow and firn cover from the top of the ice sheet and exposing older, denser ice beneath. BIAs also affect the local surface mass balance and climate (Jonsson, 1990; Orheim and Lucchitta, 1990; Bintanja and Van den Broeke, 1995; Winther and others, 1996, 2001; Liston and others, 2000; Bintanja and Reijmer, 2001; Brown and Scambos, 2004; Genthon and others, 2007; Gorodetskaya and others, 2013). These studies suggest that BIAs could be climate-sensitive. BIAs also provide optimal hunting grounds for meteorites sought for scientific studies (Koeberl, 1989; Delisle and Sievers, 1991; Maurette and others, 1991; Benoit, 1995; Harvey, 2003; Folco and others, 2006). They may also be useful in a more practical sense because of their flat, smooth and hard characteristics. BIAs have sometimes been used as landing strips for aircraft to facilitate logistics and transport (Mellor and Swithinbank, 1989; Swithinbank, 1991; Sinisalo and Moore, 2010). The use of BIAs could considerably enhance logistical support capabilities for Antarctic scientific research. Reviews of Antarctic BIAs have been written on topics including the general characteristics of blue ice, the formation and age of BIAs, glaciological, meteorological and climatological features (Bintanja, 1999). These reviews also suggest topics for further research on blue ice.

There are significant seasonal and interannual variations in BIAs closely related to seasonal changes of weather

(Brown and Scambos, 2004), so it is important to be able to identify blue ice, delineate its location and distinguish among different types of blue ice. Satellite remote sensing has greatly enhanced our knowledge of the physical properties of snow and ice. Various remotely sensed data have been employed to identify blue ice, including aerial photographs as well as images from Landsat Multispectral Scanner (MSS), Landsat Thematic Mapper (TM), Landsat Enhanced TM Plus (ETM+), Satellite Pour l'Observation de la Terre (SPOT), Moderate Resolution Imaging Spectroradiometer (MODIS), Advanced Very High Resolution Radiometer (AVHRR) and synthetic aperture radar (SAR) (Orheim and Lucchitta, 1990; Swithinbank, 1991; Bronge and Bronge, 1999; Liston and others, 2000; Winther and others, 2001; Brown and Scambos, 2004; Liu and others, 2006; Scambos and others, 2007; E and others, 2011; Yu and others, 2012). All of these studies confirmed the distributions of BIAs in various locations in Antarctica and indicated that using remotely sensed data to identify BIAs was simple, effective and relatively inexpensive. Multitemporal satellite images are needed during studies of seasonal and interannual variations of BIAs. However, it is impossible to obtain a single temporal satellite image to map BIAs throughout Antarctica given Antarctica's vast area of $\sim 14 \times 10^6 \text{ km}^2$. Nevertheless, it is important to map BIAs in Antarctica using multitemporal satellite data, which could provide information for other BIA studies. Till now, only one mapping of BIAs over the entire Antarctic continent has been implemented, using US National Oceanic and Atmospheric Administration (NOAA) AVHRR data with 1.01 km spatial resolution acquired in the 1980–94 austral summers (Winther and others, 2001). In all other studies, BIAs have been mapped only in sparse local areas. No continent-wide BIA dataset with higher spatial resolution has been assembled. The Landsat Image Mosaic of Antarctica (LIMA; Bindschadler and others, 2008) and an improved LIMA (Hui and others, 2013) now exist at 15 m spatial resolution covering the area north of 82.5° S . This paper presents a comprehensive mapping of Antarctic BIA distributions at 15 m spatial resolution using improved LIMA data covering the area north of 82.5° S (Hui and others, 2013) and at $\sim 150 \text{ m}$ resolution using the snow grain-size (SGS) image of the MODIS-based Mosaic of Antarctica (MOA) dataset covering the area south of 82.5° S (Scambos and others, 2007). This work will build the foundation for long-term monitoring of changes in Antarctic BIAs, which could be indicative of global changes, and will accumulate both data and experience for further studies.

DATA AND METHODS

ETM+ data and SGS image of MOA dataset description

This study uses 1073 scenes of Landsat-7 ETM+ images acquired during austral summers over 1999–2003 that cover the portion of the Antarctic continent north of 82.5° S , and an SGS image of the MOA dataset acquired during the 2003/04 austral summer covering the area south of 82.5° S , where Landsat has no coverage because of its near-polar orbit. The ETM+ raw data were taken for the LIMA project (Bindschadler and others, 2008) and were downloaded from the US Geological Survey website (<http://lima.usgs.gov/access.php>), and the SGS image of the MOA dataset (Haran and others, 2005; Scambos and others, 2007) was provided

by the US National Snow and Ice Data Center (NSIDC; <http://nsidc.org>).

To correct for snow and ice saturation in the visible spectrum of ETM+ data due to the high reflectance of these surfaces, a linear regression method obtained from the relationship of different bands was adopted for DN saturation adjustment of ETM+ data, and clear textures were observed after the adjustment. The Landsat ETM+ data were then converted into planetary reflectance. The multispectral bands (bands 1–5 and 7) with 30 m resolution were then sharpened into 15 m spatial resolution using band 8. Finally, the Landsat-7 ETM+ mosaic image of Antarctica north of 82.5° S was assembled by combining all multispectral ETM+ scenes at 15 m spatial resolution (Hui and others, 2013).

The SGS image of the MOA dataset was achieved from the MOA acquired during the 2003/04 austral spring and summer, and re-gridded using image super-resolution methods to 125 m spacing (150–200 m true resolution) (Scambos and others, 2007).

The ETM+ data with 15 m spatial resolution and the SGS image of the MOA dataset with 125 m gridding space were used to map the BIAs. During BIA mapping, the method of thresholds based on band ratio and reflectance was employed in ETM+ data and the method of threshold values was applied in the SGS image of the MOA dataset. Details are described in the next subsection.

Thresholds based on band ratio and reflectance for ETM+ data

Various methods have been used to distinguish BIAs from snow, firn or exposed rock, including supervised and unsupervised classification (Winther and others, 2001; E and others, 2011; Yu and others, 2012), image segmentation based on image texture (Liu and others, 2006; Yu and others, 2012), and threshold discrimination based on a band ratio or grain size obtained from a band ratio (Orheim and Lucchitta, 1990; Bronge and Bronge, 1999; Brown and Scambos, 2004; Scambos and others, 2007; E and others, 2011). Good results have been achieved using these methods. The band ratio method is based on the differences in spectral reflectance of blue ice and snow, which is a clear physical basis for discriminating between blue ice and snow, moraines or exposed rock. This method has been widely used in studies of snow grain size (Dozier, 1989; Fily and others, 1997; Nolin and Dozier, 2000; Li and others, 2001; Scambos and others, 2012). The ratio can be defined as

$$R_{ij} = \frac{\rho_i - \rho_j}{\rho_i + \rho_j}, \quad (1)$$

where ρ_i is the apparent reflectance or the ground reflectance for band i .

The ratio between two bands can remove much of the effect of illumination and minimize the influence of surface slopes. Ratios can also be used as an index for grain size, though one specific ratio cannot give an exact grain-size value. Studies show that certain ratios, i.e. those that use Landsat band 5 (1.55–1.75 μm), are more sensitive to terrain conditions than other ratios, i.e. those that use band 7 (2.08–2.35 μm). Band 4 (0.76–0.90 μm) is less sensitive to atmospheric effects and pollution than bands 2 (0.52–0.60 μm) and 3 (2.08–2.35 μm) (Dozier, 1989; Fily and others, 1997; Koren, 2009). It was found that values of R_{47} match the theoretical curves of snow grain size better than those of R_{45} , though both may have had saturation

problems in the ratio for large grain sizes in that study (Fily and others, 1997). This study selects R_{47} as the best ratio to obtain the thresholds for mapping BIAs using Landsat-7 ETM+ data. Figure 1 shows the clear differences in spectral reflectance among blue ice, snow and rocks. These differences are caused by the scattering properties of materials with different grain sizes (Dozier, 1989; Nolin and Dozier, 1993). In Figure 1 the spectral reflectance was measured in December 2009 near Zhongshan station ($69^{\circ}22'S$, $76^{\circ}22'E$), Antarctica, using an Analytical Spectral Devices (ASD) FieldSpec Pro FR spectroradiometer over the 350–2500 nm spectral range. The weather was sunny and breezy when spectra measurements were taken. The six bars in Figure 1 represent six bands (bands 1–5 and 7) of the ETM+ sensor. Figure 1 shows that snow had the highest reflectance in the visible and near-infrared (VNIR) wavelengths of the spectrum, with two peaks in reflectance in the shortwave infrared (SWIR) wavelength. Blue ice had higher reflectance than snow in the blue spectrum, hence its bluish appearance, but strongly absorbed solar radiation in the red and near-infrared portion of the spectrum and had very low reflectance in SWIR. The spectral signatures of exposed rock were very different from those of snow and blue ice. Exposed rock had strong absorption in the visible spectrum, but the reflectance changed little from the near-infrared wavelength to SWIR, and the reflectance values in SWIR were higher than those of both blue ice and snow. Comparison of the spectral curves for blue ice, snow and exposed rocks indicates that the red, infrared and SWIR bands have the most distinct spectra for discriminating blue ice from snow and exposed rocks. The slope of the reflectance curve of blue ice in band 4 was larger than that of snow in band 4, and the values of blue ice and snow in band 7 were small and relatively constant. Consequently, it was inferred that values of R_{47} for blue ice and snow should be large and that those of snow should be less than those of blue ice because of the different slopes of the reflectance curves. Values of R_{47} for exposed rocks should be negative or very small. Therefore, R_{47} can clearly distinguish BIAs from snow and exposed rocks.

To cover the whole Antarctic continent requires a great amount of Landsat-TM/ETM+ data. In this study, 14 sample ETM+ scenes in different places were randomly chosen to define the threshold values for blue-ice identification. Previous studies indicated that BIAs, snow, rock, shadowed snow or BIAs, and shadowed exposed rock could be identified and differentiated in the Landsat TM/ETM+ images. BIAs were classified as wind-induced or melt-induced, depending on the dominant climate process, which is closely related to the mechanism of formation of the BIAs (Winther, 1994; Liston and others, 1999; Winther and others, 2001). They were also classified as smooth, level or rough, according to image texture information from ETM+ data. Smooth BIAs were related to wind-induced BIAs with smooth image texture, level BIAs were associated with pond and lake ice (seasonal ice) with homogeneous texture (a level, specular surface) and rough BIAs were connected with troughs, ridges and surface crevasses with a relatively rough surface (Yu and others, 2012). Pixels of smooth BIAs and level BIAs were almost pure, but those of rough BIAs were likely to be mixed pixels containing both blue ice and snow. In this study, training areas for different surface features were identified by visual interpretation. BIAs were only classified as smooth or rough according to pixel type (pure or mixed pixels). The

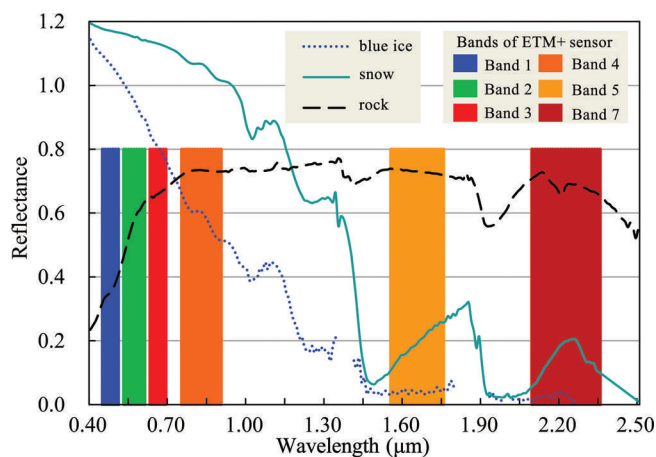


Fig. 1. Spectral reflectance of blue ice, clean snow and exposed rocks near Zhongshan station, Antarctica, in December 2009. The spectra of blue ice, snow and rocks were measured at $69^{\circ}27.242'S$, $76^{\circ}21.134'E$ on 22 December 2009, at $69^{\circ}19.314'S$, $76^{\circ}27.517'E$ on 12 December 2009 and at $69^{\circ}22.286'S$, $76^{\circ}22.061'E$ on 24 December 2009. It was sunny and breezy when the spectra measurements were taken. The six bars from left to right are the blue, green, red and near-infrared bands respectively (bands 1–4) and two SWIR bands (bands 5 and 7).

former category was closely related to small grain size, while the latter was better related to larger grain size.

Six different types of surface features were used as training samples: smooth BIAs; rough BIAs; snow; exposed rock; shadowed snow or BIAs; and shadowed exposed rock. First, we selected training samples for these six types through visual interpretation of 14 randomly selected false-color composite images, which were combined by bands 7, 4 and 1. Because the 14 sample images were randomly selected, not all the six different types of surface features were included in each sample image. Next, some statistical variables – maximum, minimum, mean, standard deviation of the apparent reflectance and R_{47} , and the number of pixels – were calculated for all the training samples in each image. Finally, comparative analyses were carried out to determine the threshold values for BIA identification.

The mean and standard deviation of apparent reflectance of six bands for each surface feature type as derived from our training samples in 14 sample images are listed in Table 1, and the mean R_{47} value of each surface feature type in each sample image as derived from our training samples is shown in Table 2.

Table 1 shows that of the six types of surface, snow had the highest apparent reflectance in the VNIR wavelengths (bands 1–4) while shadowed exposed rock had the lowest; both had only slight fluctuations and small standard deviations. The apparent reflectance of smooth and rough BIAs was the next highest after snow, but the roughest BIAs had greater ranges and high standard deviations, likely caused by the mix of snow and BIAs in the pixels. The apparent reflectance of shadowed snow and BIAs was lower, but changes in the apparent reflectance from band 1 to band 4 were successively reduced, as were those of smooth BIAs, rough BIAs, snow and shadowed exposed rock. Rock showed the highest reflectance of all six types for the SWIR wavelength (bands 5 and 7). The results are consistent with those obtained from the LIMA dataset for Lambert Glacier basin (Yu and others, 2012).

Table 1. The mean (ρ) and standard deviation (SD) of apparent reflectance of six bands for each surface feature type as derived from our training samples

Band	Smooth BIAs		Rough BIAs		Snow		Rock		Shadowed exposed rock		Shadowed snow or BIAs	
	ρ	SD	ρ	SD	ρ	SD	ρ	SD	ρ	SD	ρ	SD
1	0.842	0.042	0.862	0.331	0.955	0.023	0.226	0.100	0.170	0.023	0.390	0.100
2	0.786	0.060	0.812	0.312	0.920	0.039	0.202	0.091	0.102	0.016	0.273	0.101
3	0.686	0.084	0.752	0.293	0.904	0.032	0.199	0.092	0.064	0.013	0.203	0.109
4	0.482	0.138	0.635	0.255	0.865	0.027	0.201	0.094	0.037	0.014	0.148	0.112
5	0.019	0.014	0.034	0.017	0.082	0.014	0.179	0.083	0.011	0.008	0.013	0.013
7	0.014	0.010	0.024	0.012	0.061	0.011	0.152	0.073	0.011	0.005	0.010	0.009
Number of pixels	725 602		1 043 039		4 759 091		476 304		64 886		64 217	

Table 2 shows that R_{47} values of smooth and rough BIAs were >0.900 (minimum value 0.901), those of snow were <0.900 (maximum value 0.894), those of rock were <0.500 and those of shadowed exposed rock and shadowed snow or BIAs were unsteady with wide fluctuations and a maximum value of 0.914; the large range is likely associated with terrain, grain size and scattered light. In Table 1, the R_{47} values of smooth and rough BIAs are >0.90 , the highest for any area.

Tables 1 and 2 suggest that pixels with R_{47} values greater than 0.90 *cannot* be identified as blue ice because pixels containing shadowed snow or BIA pixels could be misidentified as blue ice. Previous studies showed that BIAs have a lower albedo (0.50–0.70) than snow (0.80–0.90) (Bintanja and Van den Broeke, 1994, 1995; Bintanja and others, 1997; Bintanja, 1999; Reijmer and others, 2001; Warren and others, 2002), so the apparent reflectance could be used as another criterion to identify BIAs. In Table 1 the reflectance of bands 3 and 4 distinguishes among the six different surface features well. Band 4 exhibits the greatest difference between itself and the other bands.

The band 4 average reflectance of smooth and rough BIAs (0.482 ± 0.138 and 0.635 ± 0.255 , respectively) was lower than that of snow (0.865 ± 0.027) and much higher than that of shadowed snow or BIAs (0.148 ± 0.112). Maximum and minimum values of reflectance of the six surface features in

the 14 sample images were obtained and plotted (Fig. 2). Figure 2 clearly shows that the maximum and minimum reflectance of both smooth and rough BIAs in band 4 were lower than those of snow and much greater than those of the other three surface features. The maximum reflectance of BIAs was lower than that of snow. Pixels with band 4 reflectance between 0.30 and 0.70 could be identified as BIAs. Therefore, the criteria to identify BIAs were defined as having both $R_{47} > 0.900$ and band 4 reflectance between 0.3 and 0.70.

The difference in spectral characteristics between smooth and rough BIAs is clearly very small (Fig. 2). It was therefore very difficult to distinguish smooth from rough BIAs using just their spectral signatures. In this study, we focused on mapping BIAs and identifying their spatial correlation with ice velocity and elevation; we gave less attention to identifying the types of BIAs.

Threshold-based snow grain size for MOA dataset

BIAs covering the area south of 82.5°S were mapped using the SGS image of the MOA dataset. BIAs were considered when snow grain size was >400 (Scambos and others, 2007).

All data were processed using a procedure developed using ENVI software (Exelis Inc., McLean, VA) and Interactive Data Language (IDL; Exelis Inc., McLean, VA, USA).

Table 2. The mean R_{47} value of each surface feature type in each sample image as derived from our training samples

ID of ETM+ image	Acquisition date	Smooth BIAs	Rough BIAs	Snow	Rock	Shadowed exposed rock	Shadowed snow or BIAs
LE7023114000131850	14 Nov 2001	0.981	0.922	0.886	0.310	0.480	0.863
LE7051117000001650	16 Jan 2000	0.961	0.925	0.853	-0.084	0.449	0.964
LE7068110000132950	25 Nov 2001	0.946	0.907	0.870	-0.002	0.781	0.973
LE7094107000235550	21 Dec 2002	0.904	0.934	0.847			
LE7110107000001450	14 Jan 2000	0.969		0.889	0.103		
LE7124108000101850	18 Jan 2001	0.974	0.972	0.859	-0.215		
LE7130112000235151	17 Dec 2002	0.955	0.935	0.876	0.028	0.305	0.876
LE7138107000232750	23 Nov 2002	0.928	0.908	0.821	0.349		
LE7151110000131950	15 Nov 2001	0.958		0.868	-0.021	0.535	0.863
LE7160122000131851	14 Nov 2001	0.976	0.901	0.894	0.167	0.751	
LE7174110000133650	2 Dec 2001	0.933	0.912	0.891	0.417		0.914
LE7173117000234850	14 Dec 2002	0.909	0.917	0.875	0.414	0.429	
LE7218109000100451	4 Jan 2001	0.932	0.937	0.855	0.267	0.715	0.853
LE7231114000302150	21 Jan 2003	0.964	0.969	0.881			
Average and SD of R_{47}		0.949 ± 0.024	0.928 ± 0.022	0.869 ± 0.020	0.144 ± 0.199	0.556 ± 0.162	0.901 ± 0.047

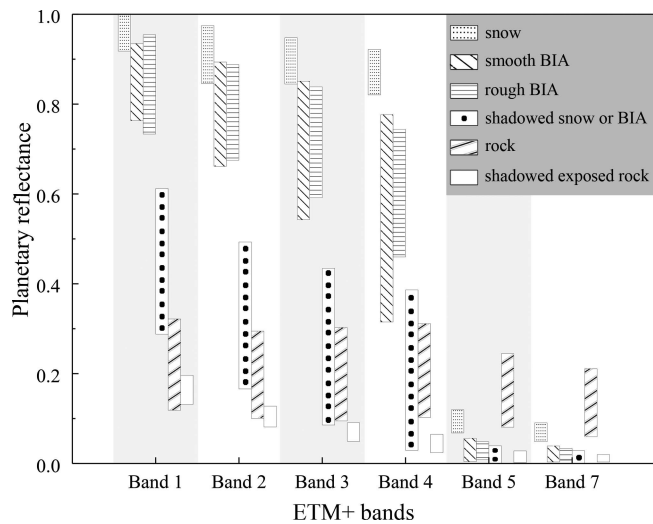


Fig. 2. Maximum and minimum reflectance values of the six surface features in the 14 sample images. In each bar the top is the maximum, and the bottom the minimum, value of apparent reflectance.

Noise reduction

BIAs were identified using R_{47} values and band 4 reflectance, and the resulting images were converted into binary images. To reduce this ‘salt-and-pepper’ effect and produce more realistic maps, a median filter with 5×5 pixel windows was applied. This filter is less sensitive than linear techniques to extreme changes in pixel values; it can remove salt-and-pepper noise and preserve useful details in images without significantly reducing the sharpness of an image. This method proved effective at removing most of the noise, though some noise still remains in the final images.

RESULTS AND DISCUSSION

We observed a total BIA of 234 549 km² across Antarctica, of which 29 402 km² was obtained from the SGS image of the MOA dataset covering the area south of 82.5° S, and the rest from ETM+ data covering the area north of 82.5° S. BIAs were usually located along the coast or near areas of exposed, ice-free rock. Figure 3 shows BIA distribution in Antarctica, and also in the Grove Mountains, where there were great differences between BIAs with salt-and-pepper noise obtained from ETM+ data and BIAs smoothed by a median filter with 5×5 pixel windows. This demonstrates the improvements achieved by the filtering process.

Accuracy assessment

It is too difficult to carry out field surveys in Antarctica to assess the accuracy of our remotely sensed BIA mapping. Instead, we use results from recently published papers as validation data. Approximately 60 000 km² and possibly as much as 241 000 km² of BIAs were observed in NOAA AVHRR data acquired during 1980–94 (Winther and others, 2001). That study’s maximum area of 241 000 km² is essentially consistent with our results of 234 549 km². The Grove Mountains contained 601.9 km² of BIAs according to ETM+ data acquired on 18 January 2001 (E and others, 2011), 610.1 km² according to MODIS data acquired on 17 January 2001 (E and others, 2011) and 745.3 km² according to MODIS data acquired during the 2003/04

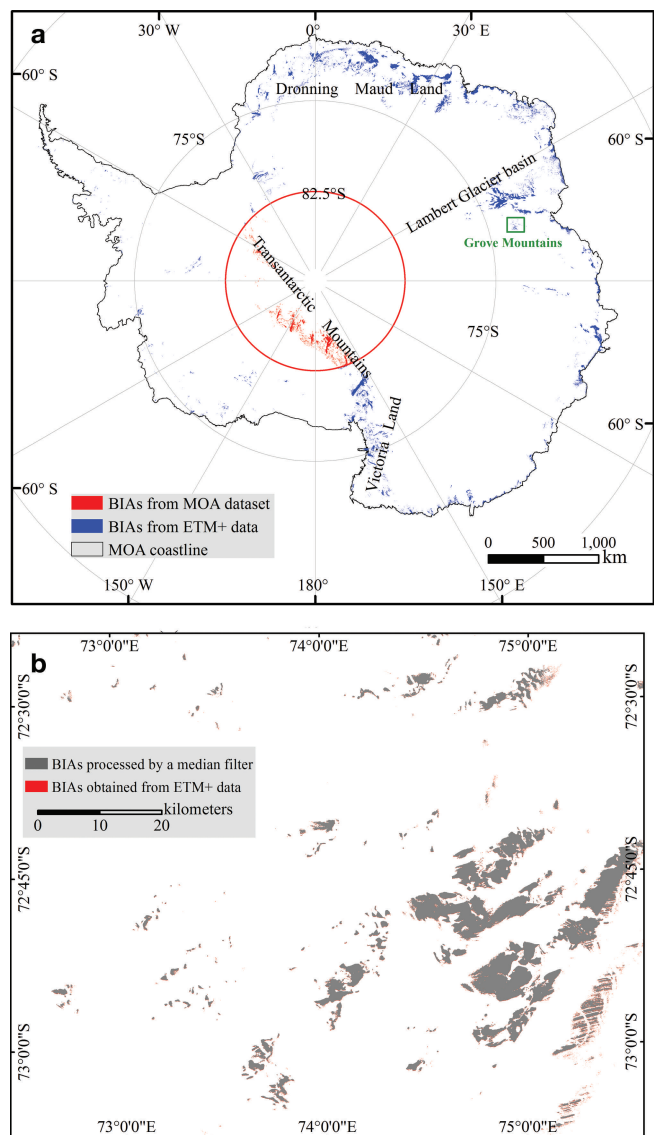


Fig. 3. The distribution of BIAs in (a) Antarctica and (b) the Grove Mountains. (a) The latitude of the red circle is 82.5° S, and within the circle (area south of 82.5° S) Landsat has no coverage because of its near-polar orbit. BIAs over the area south of 82.5° S were obtained from the SGS image of the MOA, and the others were obtained from ETM+ data. (b) Gray indicates BIAs filtered by a median filter with 5×5 pixel window size, and red indicates BIAs obtained from ETM+ data with no processing.

austral summer (Scambos and others, 2007). Our objective study found 624.2 m² of BIAs in the Grove Mountains, which agrees with the results of E and others (2011) but differs by >100 km² from the MODIS summertime data of 2003/04. Scambos and others (2007) reported five BIAs as blue-ice regions using the MOA dataset. Only one BIA was found in our result (69.566° S, 43.566° E). The other four sites were not identified, for reasons discussed in the following subsection.

The differences between previous results and our results may be due to temporal changes in BIAs, the use of different identification algorithms or the different spatial and spectral resolutions of the sensors.

The size of BIAs may have changed during the past few decades and over the years between AVHRR, MODIS and ETM+ image acquisitions due to changes in wind speed, air

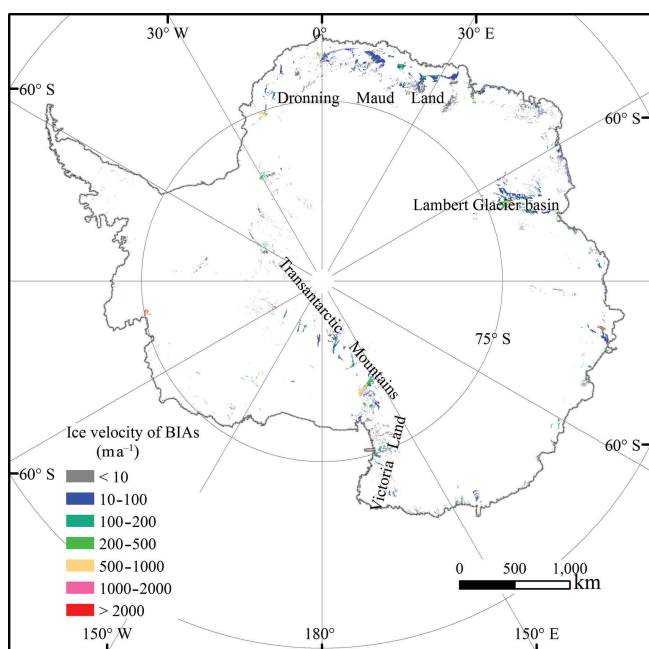


Fig. 4. BIA distribution of different classes of ice velocity.

temperature (solar radiation) and ice flow velocity, factors that could affect BIAs. The areas and spatial distributions of BIAs determined from satellite images at the same place but at different acquisition times could vary widely. Smaller areas of blue ice would become exposed when strong winds blow thin snow over BIAs, while the melting of that thin snow in periods of higher air temperatures would expose larger BIAs. The seasonal and annual changes of BIAs may be significant, so BIA changes need to be confirmed by more detailed studies.

The different spatial and spectral resolutions of the AVHRR, MODIS and ETM+ data likely also affected the differences in the various BIA calculations. Pixels containing both blue ice and snow in the 1100 m resolution AVHRR data and in the 250 m resolution MODIS data (or in the 125 m gridding-space MOA data) can either be omitted or counted as BIAs, leading to either an underestimate (pixels being omitted) or an overestimate (pixels being counted) of BIAs. However, these pixels were clearly recognized in the 15 m resolution ETM+ data. Previous studies suggest that uncertainties such as mixed pixels, crevasses and shadows could account for the difference between identified BIAs ($\sim 60\,000\text{ km}^2$) and the potential maximum area ($241\,000\text{ km}^2$) obtained from the NOAA AVHRR data (Winther and others, 2001). Figure 1 indicates that the spectral signatures of the six-band ETM+ sensor discriminated surface features more sensitively than those of the two-band AVHRR and MODIS data, which discriminated BIAs from exposed rock and snow. The higher spatial resolution and better spectral signatures of the ETM+ data improved BIA recognition. In the future, 16-bit images with nine spectral bands from the Landsat-8 Operational Land Imager (OLI) will play an important role in mapping BIAs.

The areas and spatial distribution of BIAs obtained by different methods may have differed because of different underlying assumptions and theories. In summary, our estimates for the areas and spatial distribution of BIAs are likely to be an improvement over previous results from AVHRR and MODIS data.

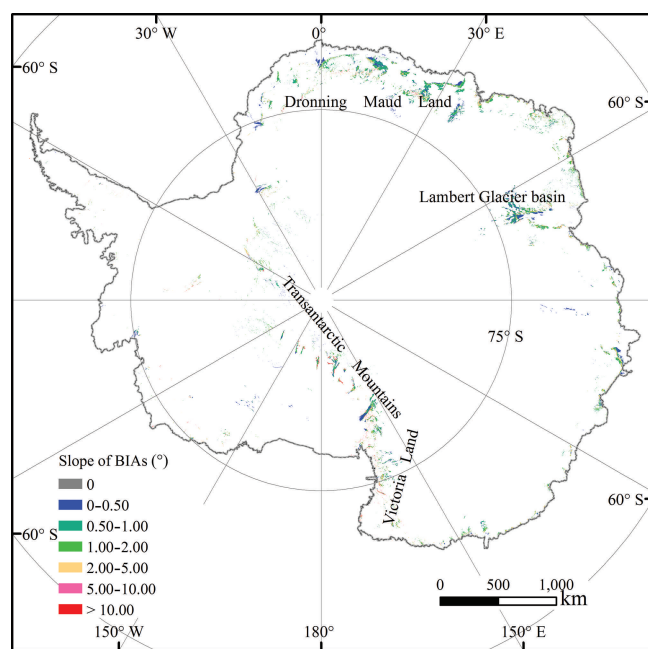


Fig. 5. BIA distribution of different classes of surface slope.

Spatial distribution

BIAs identified from 1999–2003 ETM+ data covered $\sim 1.67\%$ of the area of Antarctica and were scattered over the continent but generally located in coastal or mountainous regions. Figure 3 shows that BIAs were concentrated in four regions: Victoria Land, the Transantarctic Mountains, Dronning Maud Land and the Lambert Glacier basin. They were sparsely distributed in other regions. These results are in agreement with previous studies (Winther and others, 2001; Brown and Scambos, 2004; Scambos and others, 2007; E and others, 2011; Yu and others, 2012). These four regions either contain nunataks or mountains with high elevations and steep slopes or are at relatively low elevation and have notable glacier dynamics. BIAs seldom occur in West Antarctica and the Antarctic Peninsula. This is generally due to higher snow accumulation and a relatively brief melt season, and the predominantly springtime acquisition time of the images used in LIMA and MOA. High snow accumulation regions require an intense melt season before the winter–spring snow accumulation is converted to surface ice. Moreover, winds are more variable, and more frequently tied to synoptic weather systems rather than highly sublimating katabatic processes. Although a late-summer acquisition and BIA analysis is possible, problems would arise in distinguishing wet snow from blue ice using the methods we present here, so in our results BIAs are less distributed in West Antarctica and the Antarctic Peninsula.

BIAs were associated with surface mass balance and ice velocity features (Ligtenberg and others, 2013). Overlay analysis between BIAs and Antarctic ice velocity data with 450 m spatial resolution (Rignot and others, 2011a,b) and digital elevation models (DEMs) with 1000 m spatial resolution (Bamber and others, 2009a,b) was conducted (Figs 4 and 5). Because there were great differences in map scale between the BIA data, ice velocity data and the DEM, approximately 9342 km^2 and $15\,122\text{ km}^2$ of BIAs in the overlay analysis of ice velocity data and the DEM were ignored. Statistical analysis was performed only on overlaid data between BIAs and ice velocity and DEM.

Table 3. Areas and percentage of BIAs for different classes of ice velocity and surface slope

Ice velocity m a^{-1}	Area		Slope $^{\circ}$	Area	
	km^2	%		km^2	%
<10	81 755	36.30	0	15 138	6.90
10–100	85 483	37.96	<0.5	40 487	18.45
100–200	21 844	9.70	0.5–1	63 395	28.89
200–500	20 186	8.96	1–2	51 226	23.35
500–1000	12 653	5.62	2–5	37 082	16.90
1000–2000	2 087	0.93	5–10	6 433	2.93
>2000	1 199	0.53	>10	5 666	2.58

The results in Table 3 demonstrate that the ice velocity in ~74.26% of BIAs was $<100 \text{ m a}^{-1}$, and in about half these cases $<1 \text{ m a}^{-1}$, while 9.70% of BIAs had an ice velocity of 100–200 m a^{-1} , 8.96% an ice velocity of 200–500 m a^{-1} , and 5.62% an ice velocity of 500–1000 m a^{-1} . Ice velocity exceeded 1000 m a^{-1} for only ~1.46% of BIAs. Regions with an ice velocity of 200–500 m a^{-1} were situated on the Ross, Shackleton, Amery, Riiser-Larsen and Ronne ice shelves, which are closely associated with glacier dynamics. The results in Table 3 show that 54.24% of BIAs had surface slopes less than 1° , and 23.35% had surface slopes of $1\text{--}2^{\circ}$. Surface slopes of only 5.51% of BIAs, at most, were $>5^{\circ}$, these areas being primarily located in the Transantarctic Mountains and Victoria Land.

Surface slopes of BIAs with ice velocities greater than 500 m a^{-1} ranged from 0 to 39.05° , but 96.65% of these BIAs had slopes smaller than 5° and were distributed along ice-shelf coastlines. This suggests that BIAs with higher ice velocities may occur on flatter glaciers with smaller surface slopes. BIAs with surface slopes greater than 5° had ice velocities between 0 and 2832 m a^{-1} , but the ice velocities of 95.27% of these BIAs were $<500 \text{ m a}^{-1}$. However, BIAs with slopes greater than 5° and ice velocities greater than 500 m a^{-1} were dispersed across Antarctica. BIAs with gentle slopes but greater ice velocity were mainly distributed on coastal ice shelves, and inland BIAs with steeper slopes but slower ice velocity were mainly distributed along or near mountains. This suggests that there was no significant relationship between ice velocity and surface slope; the relationship may be evident when rough and smooth BIAs are separated.

The spatial resolution of the Antarctic ice velocity map and DEM (900 m and 1000 m, respectively) was much coarser than that of the BIA map (15 m): at least 60×60 pixels in the ETM+ data constitute only one pixel in the Antarctic ice velocity map and DEM. As a result, the ice velocity and surface slope of the BIAs could not be described at high resolution. Errors may have resulted from this large difference in spatial resolution among the datasets. The dates of data acquisition also varied greatly: ETM+ data were acquired from 1999 to 2003, ice velocity data were created from SAR data acquired from 2007 to 2009, and the DEM was generated using Satellite Radar Altimeter (SRA) data from European Remote-sensing Satellite-1 (ERS-1) acquired from 1994 to 1995, and Geosciences Laser Altimeter System (GLAS) data from the Ice, Cloud and land Elevation Satellite (ICESat) acquired from 2003 to 2008. Temporal changes could also lead to inconsistent spatial alignments of BIA locations, ice velocities and the DEM, which could also lead to errors.

CONCLUSIONS

In this study, BIAs in Antarctica have been mapped from 15 m spatial resolution Landsat-7 ETM+ images and the 125 m grid-spacing snow grain-size image of the MOA using thresholds based on band ratio and reflectance. The geographical distribution of BIAs was delineated and compared with spatial patterns in ice velocity data and surface slopes provided by DEMs. The results estimated that BIAs occupy 234 549 km^2 , or 1.67% of the continent. BIAs were scattered widely over the continent but were generally located in Victoria Land, the Transantarctic Mountains, Dronning Maud Land and the Lambert Glacier basin. The results were reasonable, realistic, and consistent with previous studies, although they differed from previous results because of possible temporal changes, different algorithms and different spatial and spectral resolutions of the sensors. The results of overlay analysis indicated that the ice velocities of ~92.92% of the BIAs were $<500 \text{ m a}^{-1}$, while 74.26% of BIAs exhibited ice velocity of $<100 \text{ m a}^{-1}$. The surface slope of most BIAs was $<5^{\circ}$, and 54.24% of BIAs had slopes of $<1^{\circ}$.

We have focused on generating a BIA dataset with high spatial resolution. This dataset will be useful for many applications, including surface mass- and energy-balance models, meteorite searches, paleoclimate studies and airstrip site selection. It can also serve as a base dataset for identifying future changes to BIAs. People will download this dataset from NSIDC soon. Temporal changes, including seasonal and annual changes, of BIAs may be significant; they may be caused by short-term snow accumulation, strong winds or other meteorological and environmental conditions. Refined results will be obtained using SAR data and new optical-sensor data (e.g. OLI data). Detailed BIA classification for applications should also be conducted in the future.

ACKNOWLEDGEMENTS

This work was supported by the Chinese Arctic and Antarctic Administration, the National Natural Science Foundation of China (grant No. 41106157), the National Basic Research Program of China (grant No. 2012CB957704), the National Natural Science Foundation of China (grant No. 41176163), the Fundamental Research Funds for the Central Universities, Chinese Polar Environment Comprehensive Investigation and Assessment Program, and the Special Foundation for Young Scientists of State Key Laboratory of Remote Sensing Science (grant No. RC12-7). Support was also provided by NASA grant NNX10AL42G. We are grateful to NSIDC for providing the snow grain-size image of the MODIS-based MOA dataset, ice velocity data and DEM data, and for hosting this

dataset and creating a web page for it. Two anonymous reviewers are thanked for their comments.

REFERENCES

- Ackert RP Jr, Mukhopadhyay S, Pollard D, DeConto RM, Putnam AE and Borns HW Jr (2011) West Antarctic Ice Sheet elevations in the Ohio Range: geologic constraints and ice sheet modeling prior to the last highstand. *Earth Planet. Sci. Lett.*, **307**(1–2), 83–93 (doi: 10.1016/j.epsl.2011.04.015)
- Bamber JL, Gomez-Dans JL and Griggs JA (2009a) *Antarctic 1 km Digital Elevation Model (DEM) from combined ERS-1 radar and ICESat laser satellite altimetry*. National Snow and Ice Data Center, Boulder, CO. Digital media: http://nsidc.org/data/docs/daac/nsidc0422_antarctic_1km_dem/index.html
- Bamber JL, Gomez-Dans JL and Griggs JA (2009b) A new 1 km digital elevation model of the Antarctic derived from combined satellite radar and laser data – Part 1: data and methods. *Cryosphere*, **3**(1), 101–111 (doi: 10.5194/tc-3-101-2009)
- Benoit PH (1995) Meteorites as surface exposure time markers on the blue ice fields of Antarctica: episodic ice flow in Victoria Land over the last 300,000 years. *Quat. Sci. Rev.*, **14**(5), 531–540 (doi: 10.1016/0277-3791(95)00011-D)
- Bindschadler R and 8 others (2008) The Landsat image mosaic of Antarctica. *Remote Sens. Environ.*, **112**(12), 4214–4226 (doi: 10.1016/j.rse.2008.07.006)
- Bintanja R (1999) On the glaciological, meteorological and climatological significance of Antarctic blue ice areas. *Rev. Geophys.*, **37**(3), 337–359 (doi: 10.1029/1999RG900007)
- Bintanja R (2000) The surface heat budget of Antarctic snow and blue ice: interpretation of temporal and spatial variability. *J. Geophys. Res.*, **105**(D19), 24 387–24 407 (doi: 10.1029/2000JD900356)
- Bintanja R and Reijmer CH (2001) Meteorological conditions over Antarctic blue-ice areas and their influence on the local surface mass balance. *J. Glaciol.*, **47**(156), 37–50 (doi: 10.3189/172756501781832557)
- Bintanja R and Van den Broeke MR (1994) Local climate, circulation and surface-energy balance of an Antarctic blue-ice area. *Ann. Glaciol.*, **20**, 160–168
- Bintanja R and Van den Broeke MR (1995) The surface energy balance of Antarctic snow and blue ice. *J. Appl. Meteorol.*, **34**(4), 902–926 (doi: 10.1175/1520-0450(1995)034<0902:TSE-BOA>2.0.CO;2)
- Bintanja R, Jonsson S and Knap WH (1997) The annual cycle of the surface energy balance of Antarctic blue ice. *J. Geophys. Res.*, **102**(D2), 1867–1881 (doi: 10.1029/96JD01801)
- Bronge LB and Bronge C (1999) Ice and snow-type classification in Vestfold Hills, East Antarctica, using Landsat-TM data and ground radiometer measurements. *Int. J. Remote Sens.*, **20**(2), 225–240 (doi: 10.1080/014311699213415)
- Brown IC and Scambos TA (2004) Satellite monitoring of blue-ice extent near Byrd Glacier, Antarctica. *Ann. Glaciol.*, **39**, 223–230 (doi: 10.3189/172756404781813871)
- Delisle G and Sievers J (1991) Sub-ice topography and meteorite finds near the Allan Hills and the Near Western Ice Field, Victoria Land, Antarctica. *J. Geophys. Res.*, **96**(E1), 15 577–15 587 (doi: 10.1029/91JE01117)
- Dozier J (1989) Spectral signature of alpine snow cover from the Landsat thematic mapper. *Remote Sens. Environ.*, **28**(1–3), 9–22 (doi: 10.1016/0034-4257(89)90101-6)
- E D, Zhang X, Wang Z and Zhou C (2011) Satellite monitoring of blue-ice extent in Grove mountains, Antarctica. *Geomat. Inform. Sci. Wuhan Univ.*, **36**(9), 1009–1011 [in Chinese with English summary]
- Fily M, Bourdelles B, Dedieu JP and Sargent C (1997) Comparison of *in situ* and Landsat Thematic Mapper derived snow grain characteristics in the Alps. *Remote Sens. Environ.*, **59**(3), 452–460 (doi: 10.1016/S0034-4257(96)00113-7)
- Fogwill CJ, Hein AS, Bentley MJ and Sugden DE (2012) Do blue-ice moraines in the Heritage Range show the West Antarctic ice sheet survived the last interglacial? *Palaeogeogr., Palaeoclimatol., Palaeoecol.*, **335–336**, 61–70 (doi: 10.1016/j.palaeo.2011.01.027)
- Folco L, Welten KC, Jull AJT, Nishiizumi K and Zeoli A (2006) Meteorites constrain the age of Antarctic ice at the Frontier Mountain blue ice field (northern Victoria Land). *Earth Planet. Sci. Lett.*, **248**(1–2), 209–216 (doi: 10.1016/j.epsl.2006.05.022)
- Genthon C, Lardeux P and Krinner G (2007) The surface accumulation and ablation of a coastal blue-ice area near Cap Prudhomme, Terre Adélie, Antarctica. *J. Glaciol.*, **53**(183), 635–645 (doi: 10.3189/002214307784409333)
- Gorodetskaya IV, Van Lipzig NPM, Van den Broeke MR, Mangold A, Boot W and Reijmer CH (2013) Meteorological regimes and accumulation patterns at Utsteinen, Dronning Maud Land, East Antarctica: analysis of two contrasting years. *J. Geophys. Res.*, **118**(D4), 1700–1715 (doi: 10.1002/jgrd.50177)
- Grinsted A, Moore JC, Spikes VB and Sinisalo A (2003) Dating Antarctic blue ice areas using a novel ice flow model. *Geophys. Res. Lett.*, **30**(19), 2005 (doi: 10.1029/2003GL017957)
- Haran T, Bohlander J, Scambos T and Fahnestock M (2005) *MODIS mosaic of Antarctica (MOA) image map*. National Snow and Ice Center, Boulder, CO. Digital media: <http://nsidc.org/data/nsidc-0280.html>
- Harvey R (2003) The origin and significance of Antarctic meteorites. *Chem. Erde – Geochem.*, **63**(2), 93–147 (doi: 10.1078/0009-2819-00031)
- Hui F and 13 others (2013) An improved Landsat Image Mosaic of Antarctica. *Sci. China: Earth Sci.*, **56**(1), 1–12 (doi: 10.1007/s11430-012-4481-5)
- Jonsson S (1990) Local climate and mass balance of a blue ice area in western Dronning Maud Land, Antarctica. *Z. Gletscherkd. Glazialgeol.*, **26**(1), 11–29
- King JC and Turner J (1997) *Antarctic meteorology and climatology*. Cambridge University Press, Cambridge
- Koeberl C (1989) Iridium enrichment in volcanic dust from blue ice fields, Antarctica, and possible relevance to the K/T boundary event. *Earth Planet. Sci. Lett.*, **92**(3–4), 317–322 (doi: 10.1016/0012-821X(89)90056-3)
- Koren H (2009) *Snow grain size from satellite images: a review*. (Note No. SAMBA/31/09) Norsk Regnesentral (Norwegian Computing Center), Oslo
- Korotkikh EV and 7 others (2011) The last interglacial as represented in the glaciochemical record from Mount Moulton Blue Ice Area, West Antarctica. *Quat. Sci. Rev.*, **30**(15–16), 1940–1947 (doi: 10.1016/j.quascirev.2011.04.020)
- Li W, Starnes K, Chen B and Xiong X (2001) Snow grain size retrieved from near-infrared radiances at multiple wavelengths. *Geophys. Res. Lett.*, **28**(9), 1699–1702 (doi: 10.1029/2000GL011641)
- Ligtenberg S, Lenaerts J, Van den Broeke M and Scambos T (2013) Blue ice areas formed by an interplay between ice velocity and SMB. *Geophys. Res. Abstr.*, **15**, EGU2013 5187
- Liston GE, Bruland O, Winther JG, Elvehøy H and Sand K (1999) Meltwater production in Antarctic blue-ice areas: sensitivity to changes in atmospheric forcing. *Polar Res.*, **18**(2), 283–290
- Liston GE, Winther JG, Bruland O, Elvehøy H, Sand K and Karlöf L (2000) Snow and blue-ice distribution patterns on the coastal Antarctic ice sheet. *Antarct. Sci.*, **12**(1), 69–79 (doi: 10.1017/S0954102000000109)
- Liu H, Wang L and Jezek KC (2006) Automated delineation of dry and melt snow zones in Antarctica using active and passive microwave observations from space. *IEEE Trans. Geosci. Remote Sens.*, **44**(8), 2152–2163 (doi: 10.1109/TGRS.2006.872132)
- Maurette M and 6 others (1991) A collection of diverse micrometeorites recovered from 100 tonnes of Antarctic blue ice. *Nature*, **351**(6321), 44–47 (doi: 10.1038/351044a0)
- Mellor M and Swithinbank C (1989) Airfields on Antarctic glacier ice. *CRREL Rep.* 89-21

- Moore JC and 7 others (2006) Interpreting ancient ice in a shallow ice core from the South Yamato (Antarctica) blue ice area using flow modeling and compositional matching to deep ice cores. *J. Geophys. Res.*, **111**(D16), D16302 (doi: 10.1029/2005JD006343)
- Nolin AW and Dozier J (1993) Estimating snow grain size using AVIRIS data. *Remote Sens. Environ.*, **44**(2–3), 231–238 (doi: 10.1016/0034-4257(93)90018-S)
- Nolin AW and Dozier J (2000) A hyperspectral method for remotely sensing the grain size of snow. *Remote Sens. Environ.*, **74**(2), 207–216 (doi: 10.1016/S0034-4257(00)00111-5)
- Orheim O and Lucchitta B (1990) Investigating climate change by digital analysis of blue ice extent on satellite images of Antarctica. *Ann. Glaciol.*, **14**, 211–215
- Rasmus K and Beckmann A (2007) The impact of global change on low-elevation blue-ice areas in Antarctica: a thermo-hydrodynamic modelling study. *Ann. Glaciol.*, **46**, 50–54 (doi: 10.3189/172756407782871774)
- Reijmer CH, Bintanja R and Greuell W (2001) Surface albedo measurements over snow and blue ice in thematic mapper bands 2 and 4 in Dronning Maud Land, Antarctica. *J. Geophys. Res.*, **106**(D9), 9661–9672 (doi: 10.1029/2000JD900718)
- Rignot E, Mouginot J and Scheuchl B (2011a) Ice flow of the Antarctic Ice Sheet. *Science*, **333**(6048), 1427–1430 (doi: 10.1126/science.1208336)
- Rignot E, Mouginot J and Scheuchl B (2011b) *MEASUREs InSAR-Based Antarctica Ice Velocity Map*. National Snow and Ice Data Center, Boulder, CO. Digital media: <http://nsidc.org/data/nsidc-0484.html>
- Scambos TA, Haran TM, Fahnestock MA, Painter TH and Bohlander J (2007) MODIS-based Mosaic of Antarctica (MOA) data sets: continent-wide surface morphology and snow grain size. *Remote Sens. Environ.*, **111**(2–3), 242–257 (doi: 10.1016/j.rse.2006.12.020)
- Scambos TA and 12 others (2012) Extent of low-accumulation ‘wind glaze’ areas on the East Antarctic plateau: implications for continental ice mass balance. *J. Glaciol.*, **58**(210), 633–647 (doi: 10.3189/2012JoG11J232)
- Sinisalo A and Moore JC (2010) Antarctic blue ice areas – towards extracting palaeoclimate information. *Antarct. Sci.*, **22**(2), 99–115 (doi: 10.1017/S0954102009990691)
- Sinisalo A, Grinsted A, Moore JC, Meijer HAJ, Martma T and Van de Wal RSW (2007) Inferences from stable water isotopes on the Holocene evolution of Scharffenbergbotnen blue-ice area, East Antarctica. *J. Glaciol.*, **53**(182), 427–434 (doi: 10.3189/002214307783258495)
- Spaulding NE and 7 others (2012) Ice motion and mass balance at the Allan Hills blue-ice area, Antarctica, with implications for paleoclimate reconstructions. *J. Glaciol.*, **58**(208), 399–406 (doi: 10.3189/2012JoG11J176)
- Swithinbank CWM (1991) Potential airfield sites in Antarctica for wheeled aircraft. *CRREL Spec. Rep.* 91-24
- Van den Broeke MR and Bintanja R (1995) Summertime atmospheric circulation in the vicinity of a blue ice area in east Queen Maud Land, Antarctica. *Bound.-Layer Meteorol.*, **72**(45), 411–438 (doi: 10.1007/BF00709002)
- Warren SG and Brandt RE (2006) Comment on ‘Snowball Earth: a thin-ice solution with flowing sea glaciers’ by David Pollard and James F. Kasting. *J. Geophys. Res.*, **111**(C9), C09016 (doi: 10.1029/2005JC003411)
- Warren SG, Brandt RE, Grenfell TC and McKay CP (2002) Snowball Earth: ice thickness on the tropical ocean. *J. Geophys. Res.*, **107**(C10), 3167 (doi: 10.1029/2001JC001123)
- Weller GE (1980) Spatial and temporal variation in the South Polar surface energy balance. *Mon. Weather Rev.*, **108**(12), 2006–2014 (doi: 10.1175/1520-0493(1980)108<2006:SATVIT>2.0.CO;2)
- Winther JG (1994) Spectral bi-directional reflectance of snow and glacier ice measured in Dronning Maud Land, Antarctica. *Ann. Glaciol.*, **20**, 1–5
- Winther J-G, Elvehøy H, Bøggild CE, Sand K and Liston G (1996) Melting, runoff and the formation of frozen lakes in a mixed snow and blue-ice field in Dronning Maud Land, Antarctica. *J. Glaciol.*, **42**(141), 271–278
- Winther J-G, Jespersen MN and Liston GE (2001) Blue-ice areas in Antarctica derived from NOAA AVHRR satellite data. *J. Glaciol.*, **47**(157), 325–334 (doi: 10.3189/172756501781832386)
- Yu J, Liu H, Wang L, Jezek KC and Heo J (2012) Blue ice areas and their topographical properties in the Lambert glacier, Amery Iceshelf system using Landsat ETM+, ICESat laser altimetry and ASTER GDEM data. *Antarct. Sci.*, **24**(1), 95–110 (doi: 10.1017/S0954102011000630)