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Curve number dependence on basic hydrologic variables governing runoff

Samuel J. Lamont
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Curve Number Dependence on Basic Hydrologic Variables Governing Runoff

Samuel J. Lamont

**Dissertation submitted to the
College of Engineering and Mineral Resources
at West Virginia University
in partial fulfillment of the requirements
for the degree of**

**Doctor of Philosophy
in
Civil and Environmental Engineering**

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2006**

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ABSTRACT

Curve Number Dependence on Basic Hydrologic Variables Governing Runoff

Samuel J. Lamont

The suitability of applying the NRCS Curve Number (CN) to continuous runoff prediction is examined by studying the dependence of the CN on several hydrologic variables. The continuous watershed model Hydrologic Simulation Program-FORTRAN (HSPF) is employed as a theoretical watershed in two numerical procedures designed to investigate the influence of soil type, soil depth, storm depth, storm distribution, and initial abstraction ratio value (λ) on the CN. This study stems from a concurrent project involving the design of a computer modeling system to support the Cumulative Hydrologic Impact Assessments (CHIA) of over 230 watersheds throughout WV. A link between the CN and HSPF soil moisture parameters is proposed for continuous runoff simulation in surface mine affected watersheds in West Virginia. A soil physics model and numerical procedure have been developed to back calculate CN's at Antecedent Runoff Condition (ARC) II from synthetic rainfall input and simulated direct runoff. A second method of CN determination is also described to provide a reference to the calculated CN values. Each HSPF parameter set, determined through calibration and by the soil physics model, is treated as a unique hypothetical watershed. It was found that the calculated CN's are highly dependent on all of the computational variables, therefore the use of the CN in continuous modeling based on antecedent soil moisture or rainfall alone does not appear to be appropriate. Differences between $\lambda = 0.05$ and $\lambda = 0.2$ are seen predominantly in the lower storm depth calculations. It is suggested that a different symbol be used to distinguish classic CN's from continuous CN's.

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Dedication

To Mom and Dad

1.0 Introduction

The most common problem encountered in hydrologic modeling is the lack of descriptive watershed data availability necessary to select the appropriate values of the model's controlling parameters. The popularity of the NRCS (Natural Resource Conservation Service) Curve Number (CN) method is based on its simplicity and embodiment of much of a given watershed's hydrologic characteristics in a single parameter. Specifically, a single parameter value embodies the ability of the surface and subsurface of the watershed to retard and capture a portion of the precipitation input, thus separating the gross precipitation into that portion that remains in the watershed, and is ultimately lost to evapotranspiration and groundwater recharge, from that portion that passes through the watershed outlet as direct runoff. As dictated by assumptions inherent in the original development of the Curve Number method, the application of the method as a model of a given watershed's separation of precipitation into losses and direct runoff requires that a single storm event be selected (of 24 hours duration, or less). Therefore, the Curve Number method is commonly termed an "event-based model", as opposed to a "continuous model". A continuous model differs from the event-based model in its ability to produce a continuous record of outflow predictions over a longer period of time, which may include many separate precipitation events occurring sequentially.

Since its original development, the Curve Number method has been modified and adapted for application in many continuous models by making additional assumptions with respect to the applicability of the basic concept to a continuously variable precipitation input. The original concept did not explicitly include time as a variable, and only predicted the total storm runoff volume. Although the basic definition of the Curve Number has not been changed, the method is now often used in continuous models in a much different context than originally intended. The typical application assumes that the Curve Number is a random variable that is some continuous function of the moisture content of the soil, in addition to the soil characteristics and hydrologic condition (average ability to infiltrate). The separation of losses from runoff is then computed continuously over time as a function of the changing CN value. This type of extended application to continuous models has provoked many questions about the validity of the assumptions required, some of which have been discussed in the literature.

The cumulative hydrologic impact assessment (CHIA) of mining on watersheds is driven by regulatory requirements that require monitoring stream water discharges and water quality parameters, and also requires a scientifically acceptable method for the prediction of mining impacts on these parameters in the future. These requirements can be addressed by application of a suitable hydrologic and water quality model. The HSPF model (Hydrologic Simulation Program Fortran, Bicknell et al. 2001) has been selected to provide these predictive estimates of mining impacts on stream water quantity/quality in the state of West Virginia. Each mine site is unique, and by definition is characterized by dramatically altered hydrologic conditions due to the extensive land disturbances. Typical changes feature altered topography, removal of vegetation and native soil structures, and highly modified drainage features that typically include drainage and sedimentation ditches, and runoff and sedimentation control detention basins. The hydrologic design of these latter structures uses the Curve Number (CN) method as an acceptable hydrologic model for the design of runoff and sediment control structures. Therefore, the CN value is generally available for any current or planned mine site, and has some legal standing due to its inclusion in various state and federal permits that are required prior to development of a new mine site. It is not the purpose of this study to address the adequacy, or lack thereof, of the Curve Number method in the design of the drainage structures, but rather to address its suitability for use in HSPF in the context of conducting the CHIA analysis.

Since the HSPF model is a complex, nonlinear, continuous model, it has many parameters that must be determined through a suitably designed calibration study on those watersheds to which it is to be applied. The application of HSPF generally involves the subdivision of the total watershed area into sub-basins, each of which can be modeled independently with regard to its rainfall inputs and corresponding outflows. If a potential mine site is to be contained within the larger watershed, it is represented as one of the many sub-basins, normally requiring that the corresponding HSPF parameters be determined through a suitable calibration procedure. The application of HSPF to hypothetical mine site sub-basin would seem to present insurmountable problems, given that calibration is not possible, and given the fact that there are no pre-existing data available to guide parameter selection. However, an intriguing possibility presents itself if one can accept that a validated Curve Number (CN) value is available. If the Curve Number can be related to those HSPF parameters that control the separation of losses from direct runoff, then perhaps the calibration requirement can be side-

stepped. This latter possibility can only be justified if the CN value is accepted as being correct, and that the method itself is accepted as appropriate for the given application. Additionally, those remaining parameters that control the other components of the watershed hydrology model must be selected via other means. In this study, it is assumed that the mine site CN value has already been validated, and that those parameters not directly related to the separation of precipitation into losses and direct runoff can be adapted from the general calibration for the whole watershed containing the mine site location(s).

The focus of this study is to investigate the possibility of use of a pre-existing mine site Curve Number (a single CN value) to select a set of surrogate HSPF parameters that govern the equivalent separation of precipitation input into losses and direct runoff. Since there is no direct method of relating a given CN value to the appropriate set of surrogate HSPF parameters, an inverse computational method is to be developed to back-calculate Curve Numbers from a set of HSPF parameters that have been derived from an intermediary soil physics model. The soil physics model serves as a sort of translator between the HSPF parameters and the Curve Number. However, this translation can not be perfect since it is not possible to equate multiple HSPF parameters to a single fixed value Curve Number. The central question to be investigated is whether or not the translation is adequate to permit the HSPF model to behave similarly to the CN method. The implications of the answer to this question reach beyond this application, to shed light on the adequacy of the use of the Curve Number in continuous hydrologic models, and further, to address a long lived controversy regarding the adequacy of the Curve Number method in general.

1.1 Objectives

As discussed above, the complexity of continuous watershed models, such as HSPF, requires the determination of a relatively large set of parameter values in order to fit the model to the hydrological characteristics of the watershed being modeled. The purpose of this research is to investigate the feasibility of using a predetermined NRCS Curve Number as guidance in the selection of those HSPF parameter values that are principal in governing the separation of precipitation into losses and direct runoff. Since multiple HSPF parameters must be related to a single CN value, it is apparent that HSPF cannot produce an exact reproduction of the separation of losses and direct runoff, under all possible hydrologic conditions, as would be produced by the

Curve Number method. This anticipated inability to produce an equivalent single valued Curve Number, across the range of variation of all input hydrologic variables, requires that a carefully designed investigation be completed to quantitatively measure the performance of HSPF in reproducing Curve Number behavior. In practice, this latter investigation will involve the completion of numerical experiments involving HSPF modeling runs that produce outputs from selected inputs, from which an equivalent CN value is calculated. Since a single valued relationship between CN values and HSPF parameters is not anticipated, there is a possibility that a limited number of functional relationships can be developed that allow the translation between the two to be practical. Before this investigation can proceed to measure this level of practicality, the procedures and algorithms that define the HSPF parameter subset, from which a corresponding CN value is computed, must be designed and developed. These research tasks will ultimately lead to several objectives being accomplished in this study:

1. Development of a translation methodology (using a suitable soil physics model) that establishes a relationship between a subset of HSPF parameters and soil characteristics that can be in turn related to CN values.
2. Completion of suitably designed numerical experiments to determine the relationship between a given subset of HSPF parameters and the Curve Number.
3. Evaluation of the practicalities of adoption of the Curve Number as a model parameter simplification technique in HSPF, and by extension, to other complex continuous watershed models.
4. Conclusions and recommendations regarding the overall applicability of the Curve Number method in HSPF applications to mined watersheds, and in watershed modeling in general.

1.2 Background

This research stems from an effort to build a computer modeling system for the prediction of the cumulative hydrologic impact assessment (CHIA) of surface coal mining on water quality and quantity in 235 Trend Station Watersheds (TSW's) in West Virginia. These TSW's were selected by the West Virginia Department of Environmental Protection (WVDEP) and are defined by water quality sampling points at their outlets. The project involves members of the Division of Resource Management and the Department of Civil and Environmental Engineering at West Virginia University, the U.S. Office of Surface Mining, and the WVDEP (Fletcher et al., 2004). The continuous watershed model HSPF was combined with the GIS-based Watershed Characterization and Modeling System (WCMS), which was developed by the Natural Resource Analysis Center (NRAC) at WVU (Strager, 2005). The CHIA modeling analysis consists of two scenarios, (1) the existing or baseline conditions and (2) the proposed surface mine site conditions.

To establish baseline conditions, a joint calibration procedure was followed using five watersheds throughout West Virginia. This resulted in one HSPF parameter set for the entire Trend Station region. Four separate watersheds were used to verify the parameter set. To model the proposed mine sites, a relationship between the Curve Number (CN) and several HSPF parameters was proposed based on a soil physics model, facilitating the use of CN's within HSPF. This relationship was developed due to the lack of runoff data needed for the calibration of HSPF to surface mine sites and because mine site CN's are available to the users of the CHIA modeling system. A numerical experiment was designed to back-calculate theoretical CN's as a function of three HSPF parameter values for a range of soil types, soil depths, storm depths, 24-hour synthetic storm distributions, and initial abstraction ratios. A second numerical experiment was designed based on the work of Hawkins (1993) to calculate CN's over the same range of watershed variables using historical precipitation records and HSPF-simulated direct runoff. In both numerical procedures, HSPF is treated as a theoretical watershed and is used to generate runoff from rainfall input.

1.3 Curve Number Method Summary

The Curve Number method for estimating direct runoff from storm rainfall was developed in 1954 by the USDA Soil Conservation Service (now known as the Natural Resources Conservation Service, NRCS). It is described in the NRCS National Engineering Handbook Section 4-Hydrology (NEH-4), Chapter 4, Storm Runoff Data (NRCS, 1993). Storm runoff depth is calculated by the expression

$$Q = \frac{(P - I_a)^2}{(P - I_a + S)} \quad \text{for } P > I_a \quad (1)$$

$$Q = 0 \quad \text{for } P \leq I_a \quad (2)$$

where Q and P are storm runoff and rainfall depths, respectively (mm), I_a is the initial abstraction, and S is the potential maximum retention when $P = I_a$. The storage index S is then transformed to the more intuitive Curve Number by the equation

$$CN = 25.4 \left(\frac{1000}{10 + S} \right) \quad (3)$$

where S is in millimeters. The Curve Number, which is dimensionless, ranges between 0 and 100 and is an index of hydrologic soil group, soil condition, land cover, and antecedent conditions. Historically, the relationship between I_a and S was fixed at $I_a = 0.2S$ where the quantity I_a / S is defined as the initial abstraction ratio (λ). Three initial watershed conditions were described by the Antecedent Moisture Condition (AMC) based on the previous five-day rainfall amount. AMC I applied to dry conditions, AMC III applied to wet conditions, and AMC II applied to the average moisture condition. It has since been recognized, however, that prior rainfall explains only part of the variation of the CN. Therefore the terminology has been changed to Antecedent Runoff Condition (ARC) (Woodward et al., 2002).

Because of its simplicity, predictability, and reliance on only one parameter, the Curve Number method has become well established in hydrologic practice with numerous applications throughout the world (Ponce and Hawkins, 1996). Typically the method is applied in one of three modes. The first and most common mode is as a frequency transform between rainfall and runoff, where a storm event of a given return period is used to predict the direct runoff corresponding to the same return period. A second mode of application is to determine infiltration rates over short time intervals for the development of flood hydrographs through use

of unit hydrographs. The third mode of application is to determine direct runoff from individual storm events imbedded in a continuous time series record. This mode is used in continuous simulation models which attempt to account for the CN variability between storm events by tracking antecedent moisture conditions through measures of previous rainfall and/or soil moisture (Hjelmfelt, et al., 2001).

Many criticisms have arisen concerning the application of the CN method since its inception. Ponce and Hawkins (1996) list several disadvantages to the method including, (1) it provides little guidance on how to vary antecedent conditions, (2) it was developed with regional data mostly from the Midwest U.S., (3) it is best suited for agricultural sites, (4) there is no accounting for spatial variability, and (5) the initial abstraction ratio is traditionally fixed at 0.2. Equation 1 can be manipulated algebraically and differentiated, provided an equation for the infiltration rate, $\frac{dF}{dt}$ (Hjelmfelt, 1980).

$$\frac{dF}{dt} = \frac{S^2}{(P - I_a) + S^2} \frac{dP}{dt} \quad (4)$$

The use of the method in this form has been criticized because of the dependence of the infiltration rate on rainfall intensity.

The method's use in continuous models has also been criticized (Van Mullem, et al., 2002). Hjelmfelt et al. (2001) state that the application of the CN method in continuous models may be completely different from the classic CN application and that more research is needed in this area. Van Mullem (1992) examined four infiltrometer studies throughout the US and found no significant relationship between soil moisture and the CN. Woodward and Plummer (2000) state that the five-day antecedent rainfall depth is not the best measure of antecedent runoff conditions and therefore it is not included in the latest version of the NEH-4 manual (NRCS, 1993).

Despite such criticisms, many variations of the CN method have been applied to continuous models. Mishra and Singh (2004) review four continuous CN models (Williams and LaSeur (1976), Hawkins (1978), Pandit and Gopalakrishnan (1996), and Mishra et al. (1998)) and propose a new variation that includes computations for the soil moisture budget, evapotranspiration (ET), surface flow routing, and baseflow contributions. Other continuous CN models include GLEAMS (Leonard, 1987), EPIC (Williams, 1987), SWAT (Arnold, 1995),

QUALHYMO (Rowney, 1992) and AnnAGNPS (Bingner and Theurer, 2001). Typically these models calculate daily ARC II CN values by defining ARC I to be the soil wilting point and ARC III to be the soil field capacity. A review of many of these and other watershed models can be found in Bora and Bera (2003).

Finally, the value of the initial abstraction ratio (λ) has also been debated. Hawkins and others (2002) studied several hundred plots of rainfall-runoff data using event analysis and model fitting to determine λ . They found that using $\lambda = 0.05$ better fit the data and is more appropriate for runoff calculations. The effect of using $\lambda = 0.05$ appeared mainly at low storm depths or lower CN values.

1.4 HSPF Model Description

HSPF is a comprehensive, continuous model designed to simulate surface and subsurface water quantity and quality processes occurring in a watershed. Its origins can be traced to the Stanford Watershed Model which was developed in the 1970's. Today, HSPF is supported by the EPA (2000). It has over twenty parameters defined in its User's Control Input (UCI) file (Bicknell 2001), many of which must be determined through calibration. Surface and subsurface flow drains from pervious land use/cover categories, (PERLND's), which are assigned unique sets of model parameters, into the appropriate stream segments (RCHRES's). Figure 1 is a schematic of the PERLND module describing its various storages and parameters.

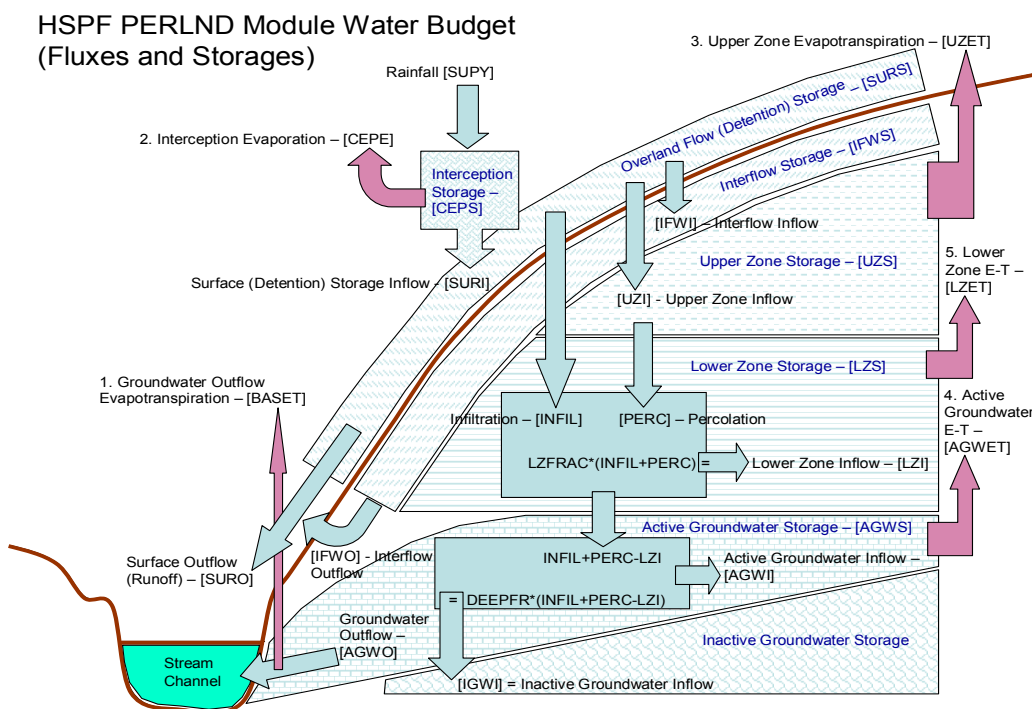


Figure 1. Flow Schematic and Storage Components within the HSPF PERLND Module

The minimum model inputs are precipitation and potential evapotranspiration (PET) time series while each of the computed storages and fluxes can be output in time series format. HSPF has been applied to a large number of watershed studies in a wide variety of locations. Forty-five studies using the model in the United States have been summarized in a user-friendly software package called HSPFParm (Donigian et al., 1999). Sams and Witt (1995) calibrated HSPF to two surface mined watersheds in Fayette County, PA, providing local relevance to this study.

1.5 Baseline HSPF Calibration Summary

The WVDEP Trend Station Watershed water quality sampling points rarely coincided with USGS stream gaging stations required for model calibration. This fact, along with the obvious impracticality of individually calibrating to 235 watersheds, led to the adoption of a joint-calibration strategy following the work of Donigian (2002) and Dinicola (1990, 2001). Five calibration watersheds scattered throughout the state were selected with the intent of finding one parameter set for all of the trend station watersheds (Figure 2). Five additional verification watersheds were used to test the validity of transferring the resulting parameters (Figure 3). The Big Sandy watershed was used for both calibration and verification by using different simulation

time periods. This resulted in one parameter set for each of the nine land use categories for the entire trend station region. The land use categories (Forest, Pasture/Grassland, Urban/Developed, Existing Mine Land, Barren Land, Shrubland, Row Crop Agriculture, Surface Water, and Wetland) were based on 1993 GAP data (Strager and Yuill, 2002).

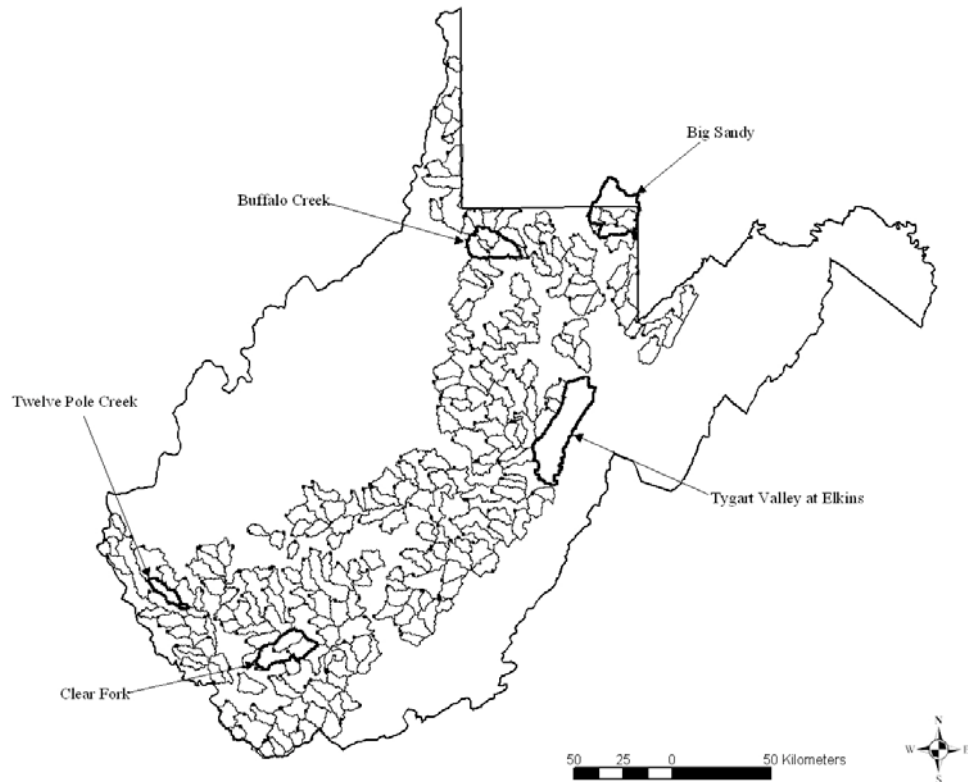


Figure 2. West Virginia CHIA Trend Stations and Calibration Watersheds.

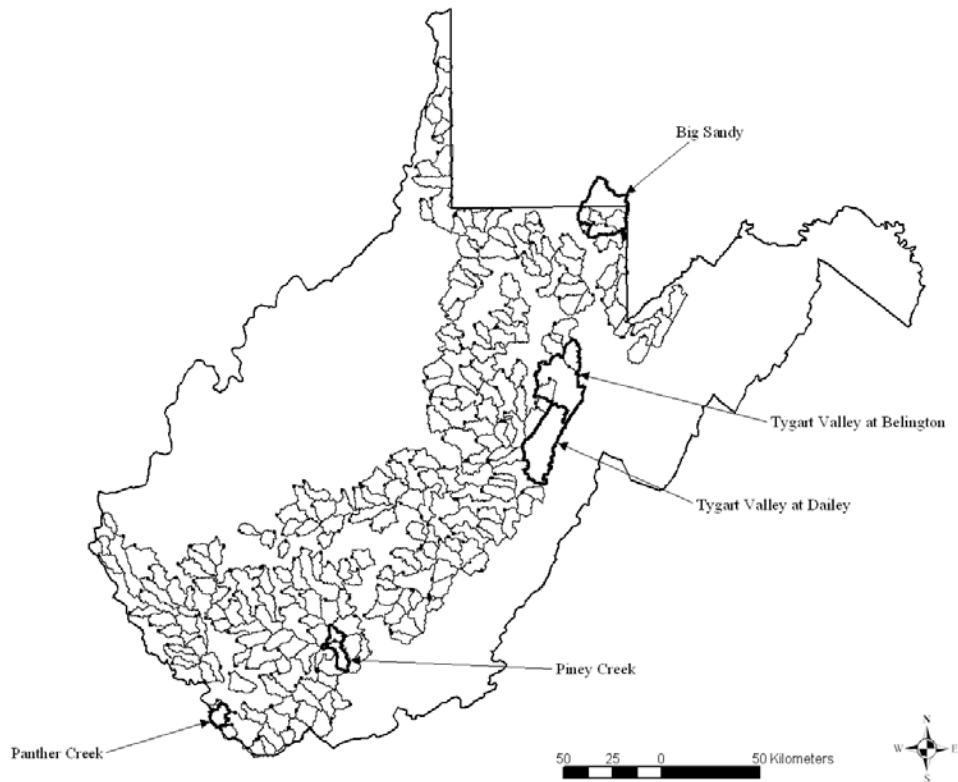


Figure 3. West Virginia CHIA Trend Stations and Verification Watersheds

The joint-calibration procedure involved two approaches. The first used the USGS semi-automated software HSPEXP (1994) which provides statistical and graphical error measures as well as parameter adjustment advice. A second calibration study was conducted using the independent parameter optimization package, PEST (Doherty, 2002). The final parameter set was selected through comparison of seven performance evaluating indices including the Coefficient of Determination (r^2), the Coefficient of Efficiency (E), and the Root Mean Square Error ($RMSE$). The simulated mean error is less than 12% for calibration watersheds and less than 15 % for four of the verification watersheds. An unpublished technical report describing the calibration procedure in detail is included as Appendix A. Several agencies and individuals contributed to this report including the West Virginia Department of Environmental Protection, the U.S. Office of Surface Mining, Kate Flynn of the USGS, Reston, VA, Jim Sams of the USGS, Pittsburgh, PA, Dr. Robert Eli of the Department of Civil and Environmental Engineering at West Virginia University, and Elena Hoeg of the Natural Resource Analysis Center at WVU. The UCI file containing the final calibration parameter set is included as Appendix B.

1.6 Soil Physics Model

1.6.1 Relating HSPF Parameters to NRCS CN using a Soil Water Physics Model

An analytical link was established between HSPF soil moisture parameters and physical soil attributes by adopting a soil water physics model based on the Green-Ampt (Green and Ampt, 1911) and the Brooks-Corey (Brooks and Corey, 1964) equations. The soil model is described in terms of the soil pore size distribution index (λ_{ps}), soil porosity (η), soil water suction head (ψ), and soil moisture content (θ) shown in Figure 4.

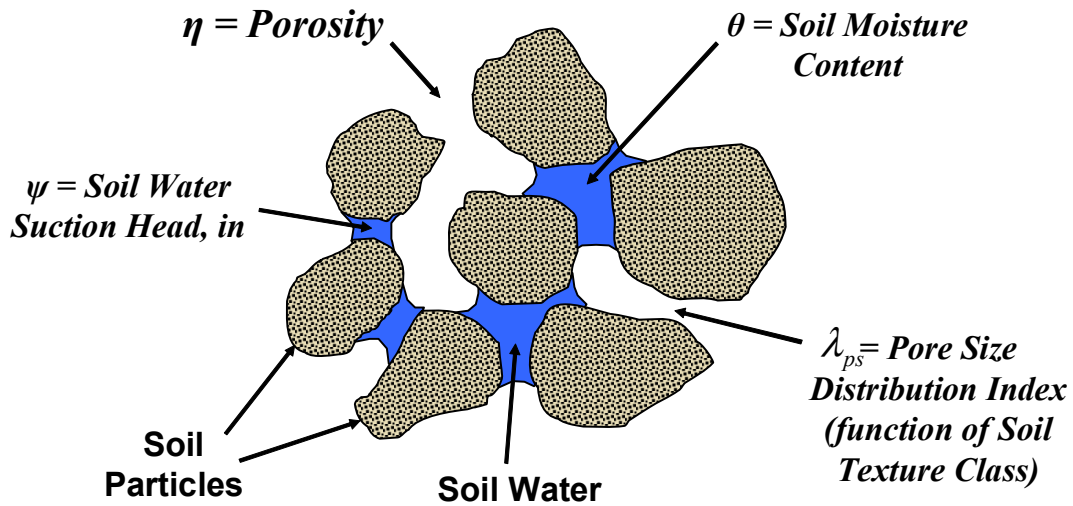


Figure 4. Soil Microstructure and Soil Water Variables (Soil Physics Model)

Brooks and Corey (1964) developed an empirical relationship between soil water suction head ψ (cm of water) and effective saturation s_e , as a function of soil texture. The Brooks-Corey equation is

$$s_e = \left[\frac{\psi_b}{\psi} \right]^{\lambda_{ps}} \quad (4)$$

where ψ_b is the soil water suction head at which air first enters the soil (called the bubbling pressure) and λ_{ps} is pore size distribution index (a function of soil texture). The effective saturation is defined by

$$s_e = \frac{\theta - \theta_r}{\eta - \theta_r} \quad (5)$$

where θ is the moisture content of the soil (cm^3/cm^3), θ_r is the residual moisture content of the soil (equivalent to the wilting point), and η is the soil porosity (see Figure 5 below).

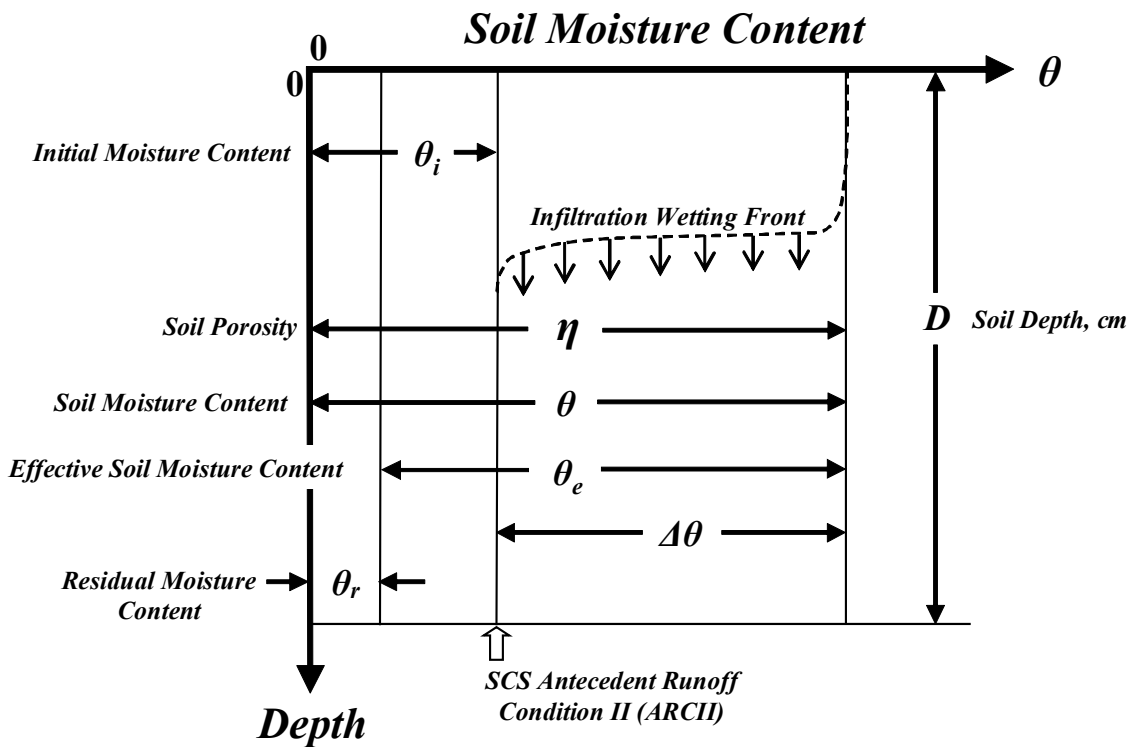


Figure 5. Soil Moisture Content as a Function of Soil Depth (Soil Physics Model)

Referring to Figure 5, the idealized soil water physics model assumes that the soil has homogeneous characteristics over the soil depth D . Neither the HSPF PERLND nor the CN method assumes that the soil has an explicit depth. In the soil water physics model, depth is required in order to compute soil water storage depth; and therefore, the soil depth is considered

to be the “equivalent soil depth” that produces the desired storage capacity of the soil. It should be noted that the maximum possible soil moisture content is equal to the porosity η . The actual maximum soil moisture content will be slightly less than the porosity since a small amount of trapped air will remain in the soil when is fully saturated. In the development that follows, the moisture content at saturation will be assumed to be equal to the porosity since the simplification introduces a negligible error. The effective moisture content, θ_e , is the amount of moisture that can be removed by gravity drainage and plant transpiration, assuming that the soil is initially saturated. The NRCS antecedent runoff condition (ARC I, ARC II, ARC III, or an intermediate value) is determined by initial moisture content, θ_i , present in the soil prior to a storm event. To simplify the model description, the moisture content is assumed to be constant over the soil depth at any given point in time.

Brakensiek, Engleman, and Rawls (1981) used the Brooks-Corey equation (4) to develop a method to determine parameters for the Green-Ampt infiltration equation (1911). The Green-Ampt equation is

$$f(t) = K \left[\frac{\psi \Delta \theta}{F(t)} + 1 \right] \quad (6)$$

where $f(t)$ is the infiltration capacity at time t , K is the unsaturated hydraulic conductivity, ψ is the wetting front capillary pressure head, $\Delta \theta$ is the change in soil moisture content across the wetting front (Figure 5), and $F(t)$ is the accumulated infiltration at time t . Rawls, Brakensiek, and Miller (1983) used this same method to analyze approximately 5000 soil horizons across the United States to determine average values of the Green-Ampt parameters for different soil texture classifications. Table 1 lists 11 soil texture classifications used in this latter study, ranging from Sand (coarse particles) to Clay (very fine particles). Combining equations 5 and 6, and solving for θ yields

$$\theta = \theta_r + (\eta - \theta_r) \left[\frac{\psi_b}{\psi} \right]^{\lambda_{ps}} \quad (7)$$

which relates soil moisture content θ to soil water suction head ψ for a particular soil texture classification (for constant values of η , θ_r , ψ_b , and λ_{ps}). Equation 7 permits computation of the initial moisture content of the soil, θ_i , for any desired antecedent runoff condition (ARC) prior to a given storm event.

1.6.2 Computation of Equivalent HSPF Parameters for NRCS Curve Numbers

Examination of the HSPF PERLND module algorithms identifies six of the 20 parameters that have principal influence on the infiltration and soil moisture storage processes and the shape of the direct runoff hydrograph:

$UZSN$ = Upper zone nominal soil moisture storage (mm).

$LZSN$ = Lower zone nominal soil moisture storage (mm).

$INFILT$ = Index to the mean infiltration rate (mm/hr).

$INFEXP$ = Infiltration exponent parameter.

$INTFW$ = Interflow inflow parameter.

IRC = Interflow recession parameter (1/day).

The first four parameters predominate in the control of the infiltration and soil moisture storage processes, and the last two parameters predominate in the control of the shape of the direct runoff hydrograph. The HSPF model has two soil water storage variables, the upper zone storage UZS (mm) and the lower zone storage LZS (mm) (see Figure 1). The corresponding nominal storage capacities, $UZSN$ and $LZSN$ (mm) are user adjustable model fitting parameters that are a function of “precipitation patterns and soil characteristics”, according to BASINS Technical Note 6 (2000). The application of these nominal storage capacities in HSPF algorithms (Bicknell, et al, 2001) implies the following relationship between the nominal storages and the effective maximum storage capacities:

$$\begin{aligned} UZS_{\max} &= 3.0(UZSN) \\ LZS_{\max} &= 2.5(LZSN) \end{aligned} \quad (8)$$

In view of the PERLND model component design, as shown in Figure 1, there is no defined soil depth and the combined values of UZS and LZS are the total of all storage in the subsurface between the soil surface and the ground water table (neglecting the short term interflow storage). The description of the function of the upper zone storage UZS , as stated in Hydrocomp (1969) (original source of the PERLND algorithm), is to provide for “depression storage and storage in highly permeable surface soils”. It is further stated that “the upper zone storage prevents overland flow from a portion of the watershed depending on the value of the ratio $UZS/UZSN$, but since the nominal capacity $UZSN$ is small, the upper zone retention percentage decreases rapidly with early increments of (rainfall) accretion”. Inflow to the upper zone is governed by the storage ratio $UZS/UZSN$ alone and is not considered to be part of the

infiltration process (Bicknell, et al, 2001). In view of these latter interpretations, the equivalent soil depth D is assumed to be defined by the maximum effective storage capacity of the lower zone storage, LZS_{max} :

$$D = \frac{LZS_{max}}{\eta - \theta_r} \quad (9)$$

As already noted above, η is the soil porosity and θ_r is the residual moisture content. Combining equations 8 and 9 produces:

$$LZSN = \frac{(\eta - \theta_r)D}{2.5} \quad (10)$$

Donigian and Davis (1978) presented guidelines on the ratio of the nominal capacities of the two storages, $UZSN/LZSN$. They recommended that the nominal storage capacity of the upper zone $UZSN$ be from 0.06 to 0.14 of that for the lower zone $LZSN$. Therefore, an average ratio of 0.10 was selected:

$$\frac{UZSN}{LZSN} = 0.1 \quad (11)$$

Combining equations 10 and 11, and solving for $UZSN$, yields:

$$UZSN = \frac{(\eta - \theta_r)D}{25} \quad (12)$$

The antecedent soil water depth of the lower zone storage, LZS_i , corresponding to the NRCS type II antecedent runoff condition (ARC II) (SCS, 1986) can be computed for the effective soil depth D if the corresponding soil moisture content θ_i is known:

$$LZS_i = (\theta_i - \theta_r)D \quad (13)$$

Rawls and Brakensiek (1986) conducted studies comparing the runoff volume predictions of the Green-Ampt infiltration model to the CN model. They concluded that $\psi = 340$ cm was equivalent to the NRCS antecedent runoff condition II (ARC II). Using this value in equation 7 for each soil texture class results in the initial soil moisture content value θ_i , which in turn can be used to compute the antecedent soil water depth using equation 13. Table 1 lists values of θ_r , η , λ_{ps} , ψ_b , and θ_i , for each soil texture class.

The remaining parameters required to establish the HSPF and CN relationship for the design storm direct runoff volume are the infiltration parameter $INFILT$ (mm/hr) and the

infiltration exponent parameter *INFEXP*. The HSPF infiltration capacity *IBAR* (mm/hr) is computed by (Bicknell, et al, 2001)

$$IBAR = \left[\frac{INFILT}{\left(\frac{LZS}{LZSN} \right)^{INFEXP}} \right] INFFAC \quad (14)$$

where *INFFAC* is the frozen ground adjustment factor (set to 1 for unfrozen ground) and *INFEXP* is set equal to 2, consistent with typical applications of HSPF (U.S. EPA, 2000), and as recommended by Hydrocomp (1969). The values of *INFILT* for each of the soil texture classes listed in Table 2 are consistent with those values of *INFILT* recommended by BASINS Technical Note 6 (U.S. EPA, 2000) for NRCS Hydrologic Soil Groups A, B, C, and D, as listed in Table 2. The soil texture classes in Table 1 were first classified by hydrologic soil group using the soils data published by Nearing, et al., (1996). They compared NRCS Curve Numbers and hydrologic soil group classification to Green-Ampt hydraulic conductivities for a large number of soils covering a complete range of soil texture classes; therefore, it was possible to assign the proper hydrologic soil group to the soil texture classes in Table 1, according to soil texture class description and Green-Ampt hydraulic conductivity. After the appropriate hydrologic soil group

classifications were determined, values of *INFILT* from Table 2 were assigned to each soil texture class so that the values varied smoothly from Sand to Clay, and so that the boundaries between hydrologic soil group classifications reflected the limits on the range of *INFILT* values listed in Table 1. In practice, this was accomplished by plotting estimated values of *INFILT* versus Green-Ampt infiltration capacity (at $F(t) = 1$ cm), and then adjusting the *INFILT* values by trial until a smooth curve fit was achieved (see Figure 6). The remaining less critical HSPF parameters were fixed at the values determined by the calibration process (Appendix A) and with guidance from Sams and Witt (1995) for the surface mine land cover condition. Therefore, each soil texture class (Table 1) represents a surface mine site land cover condition with a unique infiltration capacity.

Table 1. Soil Texture Class Hydraulic Properties

Soil Texture Class	Total Porosity η	Residual Moisture Content θ_r	Pore Size Distribution Index λ	Bubbling Pressure ψ_b (cm)	Initial Soil Moisture Content θ_i at AMCII	Green-Ampt Soil Hydraulic Conductivity K (cm/hr)	Wetting Front Cap. Pressure Head ψ (cm)	Green - Ampt $\Delta\theta$	Green-Ampt Infiltration Capacity at $F(t) = 1$ cm $f(t)$ (cm/hr)	Hydro-logic Soil Group ¹	INFILT Estimate ² (cm/hr)
Sand	0.437	0.020	0.546	17.340	0.102	11.78	4.95	0.335	31.309	A	2.50
Loamy Sand	0.437	0.036	0.449	9.078	0.117	2.99	6.13	0.320	8.850	A	1.45
Sandy Loam	0.453	0.041	0.378	16.777	0.173	1.09	11.01	0.280	4.450	A	1.00
Silt Loam	0.501	0.015	0.207	43.337	0.332	0.65	16.68	0.169	2.479	B	0.65
Loam	0.463	0.029	0.246	23.196	0.253	0.34	8.89	0.210	0.974	B	0.27
Sandy Clay Loam	0.398	0.068	0.345	25.868	0.204	0.15	21.85	0.194	0.786	C	0.24
Silty Clay Loam	0.471	0.039	0.164	36.855	0.339	0.10	27.30	0.132	0.460	C	0.19
Clay Loam	0.464	0.155	0.259	27.249	0.316	0.10	20.88	0.148	0.410	C	0.18
Silty Clay	0.479	0.056	0.186	27.167	0.320	0.05	29.22	0.158	0.282	C	0.14
Sandy Clay	0.430	0.109	*	*	0.277 ³	0.06	23.90	0.153	0.279	C	0.13
Clay	0.475	0.090	0.187	32.917	0.339	0.03	31.63	0.136	0.159	D	0.05

All values derived from Brakensiek, Engleman, and Rawls (1981) and Rawls, Brakensiek, and Miller (1983), unless otherwise noted.

[1] Nearing, Liu, Risse, and Zhang (1996). [2] Figure 5. [3] Estimated using Rosetta (1999). [*] Unavailable.

Table 2. INFILT versus Hydrologic Soil Group (BASINS Technical Note 6, U.S. EPA, 2001)

SCS Hydrologic Soil Group	INFILT Estimate (mm/hr)	Runoff Potential
A	25.4 – 63.5	Low
B	6.35 – 25.4	Moderate
C	3.175 – 6.35	Moderate to High
D	0.635– 3.175	High

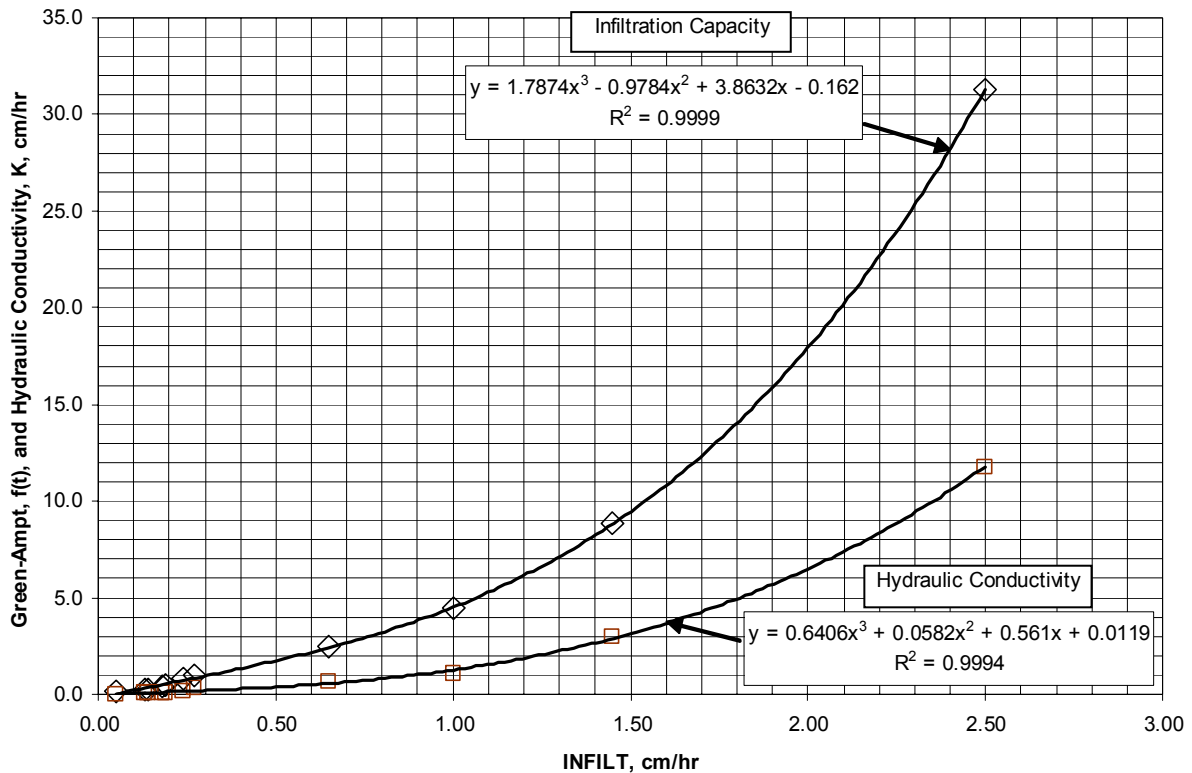


Figure 6. INFILT as a Function of Green-Ampt Infiltration Capacity and Soil Hydraulic Conductivity

2.0 Methodology

2.1 Curve Number Computation Using Cyclic Storm Input

Curve Numbers were determined numerically through an iterative process using HSPF with synthetic rainfall and potential evapotranspiration input in hourly increments

for the range of variables shown in Table 3. The HSPF parameters (LZSN, UZSN, and INFILT) and the simulated soil moisture content corresponding to ARC II (LZS_i) were calculated for each equivalent soil depth using equations 10, 11, and 13 and Table 1. Therefore, each soil depth corresponds to a theoretical watershed with unique hydrologic characteristics. The remaining HSPF parameter values were determined through the joint calibration procedure and were fixed throughout this study (Appendix B). The input rainfall time series consisted of repetitive, regularly-spaced twenty-four hour storm events of constant distribution and depth (a cyclic storm input). Four synthetic storm distributions were used to examine their possible affect on the calculated Curve Numbers (Figure 7). It should be noted that the WDM (Watershed Data Management) Triangular distribution is used for the disaggregation of rainfall records in the EPA's software for managing meteorological time series data, WDMUtil (Hummel et al., 2001).

Table 3. CN Computational Variables

Soil Type	Soil Depth (cm)	Storm Depth (mm)	Storm Distribution	λ
Sand	10	10	NRCS Type II	0.2
Loamy Sand	15	20	Uniform	0.05
Sandy Loam	20	30	Full Triangular	
Silt Loam	25	40	WDM Triangular	
Loam	30	50		
Sandy Clay Loam	35	60		
Silty Clay Loam	40	70		
Clay Loam	45	90		
Silty Clay	50	110		
Sandy Clay	60	130		
Clay	70	150		
	80			
	90			
	100			
	120			
	140			
	160			
	180			
	200			

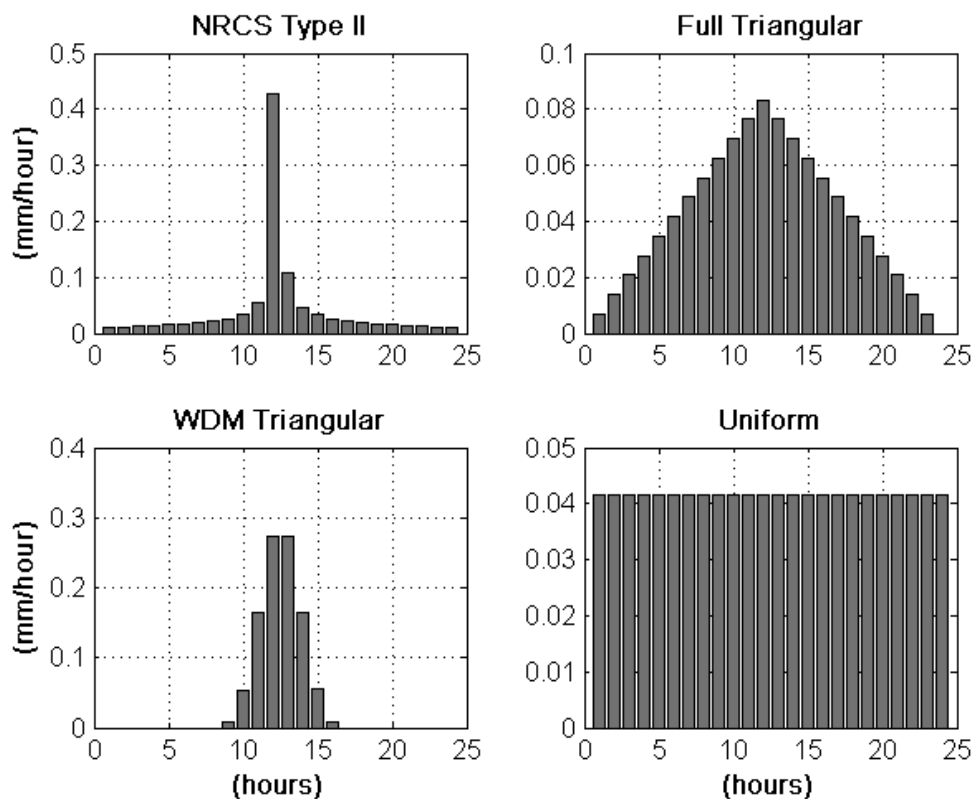


Figure 7. Synthetic Storm Distributions, Hourly Time Increment, 1 mm Accumulated Depth

The potential evapotranspiration (PET) time series consisted of a uniform rate maintained at a fixed value for all simulations consistent with a typical dry day rate (mm/hr) observed during the growing season (Figure 8). The PET was set to zero during the storm event. No diurnal fluctuation was used since the only purpose of the PET was to draw the soil moisture level down to the ARC II condition prior to the next cyclic storm event, and it was desirable not to introduce any unnecessary fluctuations into the simulation. Each simulation run was conducted over a sufficient number of storm cycles to ensure that cyclic equilibrium was reached in all of the HSPF PERLND output time series variables. Samples of the cyclic HSPF input and output are shown in Figure 8.

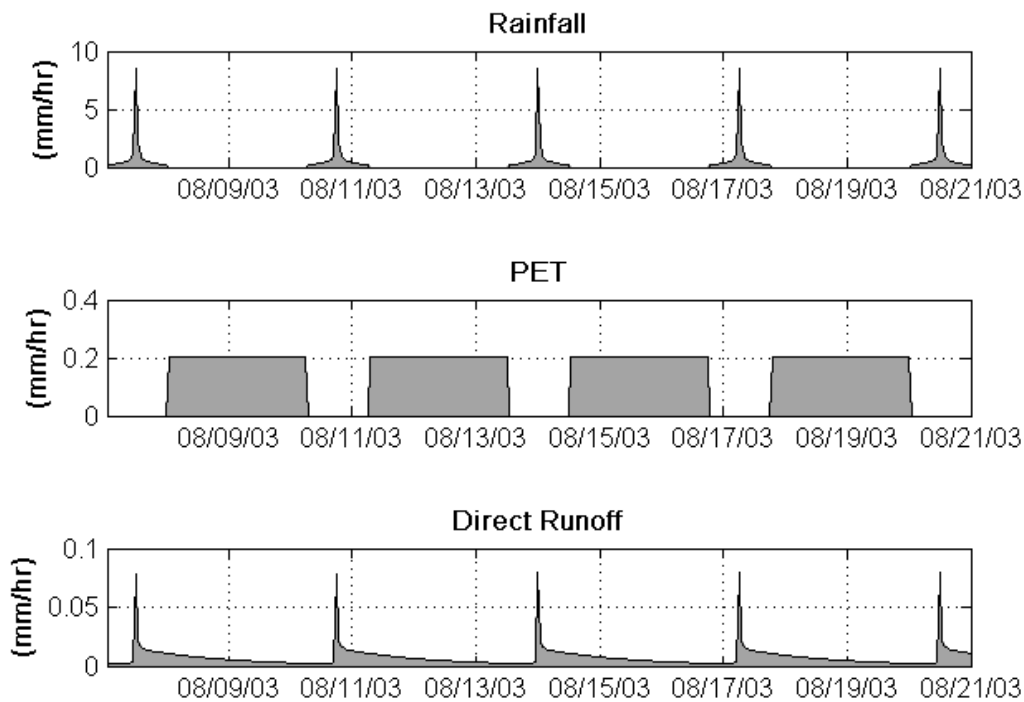


Figure 8. Sample Cyclic HSPF Input (Type II Rainfall and PET) and Output (DRO)

The time between storm events was varied by trial until the lower zone storage, *LZS*, matched the initial ARC II condition computed using equation 13. The CN was then computed using equations 1-3 with the known value of *P* (corresponding to a specific storm distribution) and the numerically determined value of *Q* (the sum of *SURO* and *IFWO* HSPF output components between storm events). A check was included to ensure the rainfall depth satisfied the condition of equation 1. This procedure was performed over the ranges of each variable listed in Table 3 for the Silt Loam and Clay Loam soil types. These soil types were selected because of their relatively high and low infiltration rates, respectively (Table 2). Additionally, the computation time required to include all eleven soil types was considered prohibitive for this study. Each soil texture class requires 836 individual HSPF simulations in order to complete all combinations of the variables listed in Table 3. This excludes the Asymptotic Method (presented later) as well as post-processing requirements.

2.2 Asymptotic Method Using Observed Storm Input

A second method of determining CN values from time series rainfall-runoff data was introduced to provide a reference to the values calculated in the synthetic cyclic storm procedure based on the soil physics model. This method follows Hawkins' (1993) asymptotic method of determining CN values from data for individual watersheds which has become widely recognized (VanMullem, et al., 2002; Hjelmfelt, 2001). It is based on the idea that the CN method is best suited to frequency transform applications. Each storm and its corresponding runoff events are extracted from a single time series record. These data pairs are then sorted individually by depth from high to low and are re-paired, ensuring that each rainfall-runoff pair are of equal return periods, even though they may not coincide in time. When the CN's are calculated from these ordered pairs and are plotted on the y-axis against storm depth on the x-axis, three trends often emerge. The first is known as complacent behavior where the CN decreases steadily with storm depth without approaching a constant value (Figure 9). CN's cannot safely be determined from data which exhibit this pattern because no constant value is approached. This trend has been found to indicate a partial source area situation (Hawkins, 1979; Pankey and Hawkins, 1981).

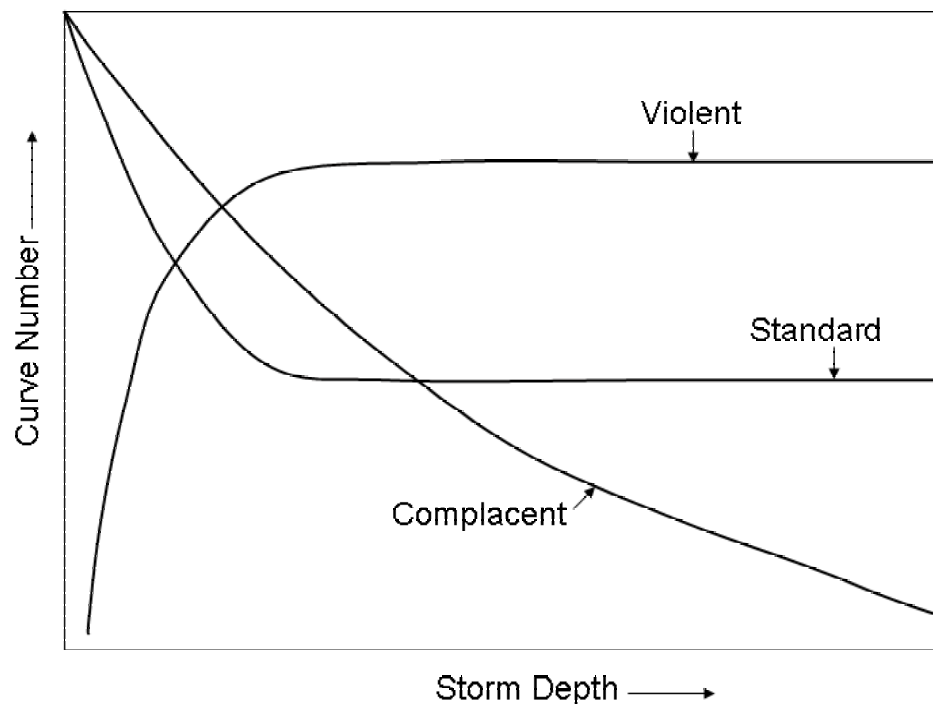


Figure 9. Sketch of Behavioral Trends

The second trend is referred to as standard behavior, where the CN values decrease with increasing storm depth and approach a constant value (Figure 9). Hawkins found this to be the most common scenario, and hypothesized that runoff generation may include a variety of processes such as overland flow and interflow. He found that equation 15 can be used to fit the Standard behavior CN-P data sets, where k is a fitting constant and CN_{∞} is the value that is approached as P increases. The value of CN_{∞} is used as the Curve Number identified with an individual watershed (Hawkins, 1993).

$$CN(P) = CN_{\infty} + (100 - CN_{\infty})\exp(-kP) \quad (15)$$

The third variation is known as violent behavior where the CN's rise suddenly with rainfall and then asymptotically approach a constant value. Violent behavior could indicate a threshold phenomenon at some critical rainfall depth. Hawkins also found that violent behavior was often accompanied by complacent behavior at lower rainfall depths. In this latter case, equation 16 (Hawkins, 1993) has been found to represent the CN-P data sets, ignoring any complacent behavior.

$$CN(P) = CN_{\infty}[1 - \exp(-kP)] \quad (16)$$

It should be noted that Figure 9 is only a sketch showing the characteristic curve shapes of the three behavior types. It does not necessarily indicate relative positioning.

Hawkins' asymptotic method of CN determination was adopted for this study using simulated direct runoff from HSPF and long (a minimum of 20 years) historical precipitation records as model input. The precipitation records were gathered from four National Climatic Data Center (NCDC) stations scattered throughout the coal mining region in WV (Figure 10).

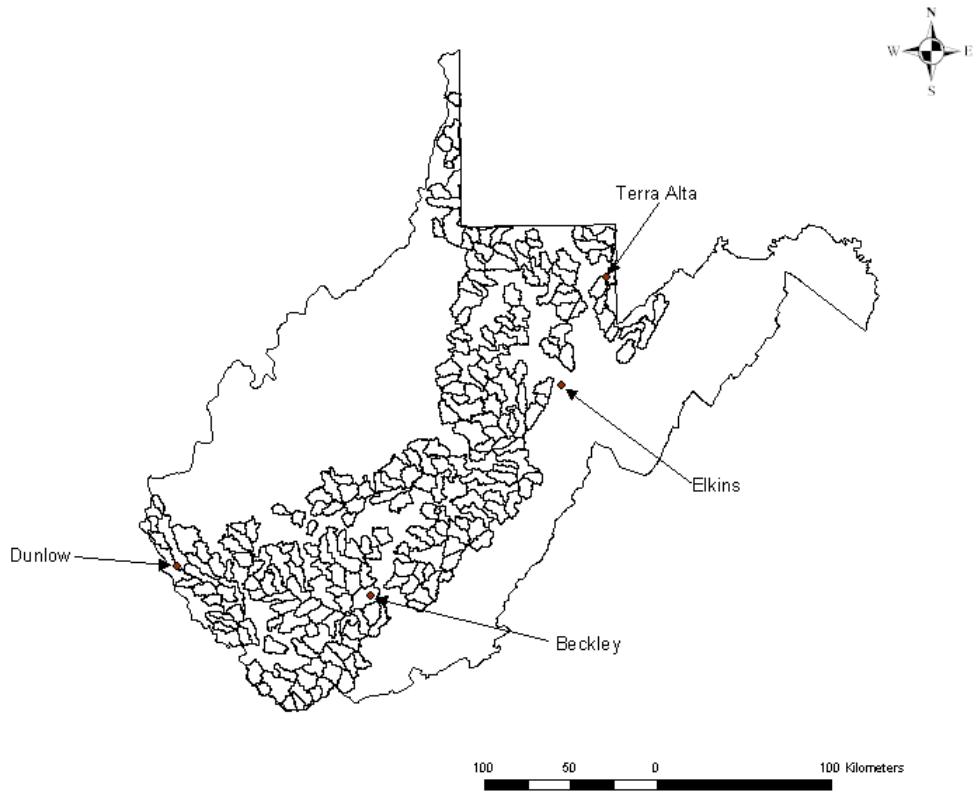


Figure 10. Location of Rainfall Gages Used in the Asymptotic Method

PET was calculated from daily minimum and maximum temperature records at the same gages and latitude using the Hamon method in the EPA’s WDMUtil software package (Hummel et al., 2001).

An algorithm was written that automatically selected the input rainfall and simulated hourly runoff events based on the following criteria. A storm event must fall within a twenty-four hour window with zero total rainfall during the previous and subsequent 24 hour periods. Additionally, the total storm depth must be greater than a designated minimum depth; in this case 13 mm (about 0.5 inches). The condition specified by equation 1 was also enforced here. The storm event search was limited to the beginning of May to the end of September to exclude snowfall. The corresponding simulated direct runoff for each event was accumulated between the first hour of the selected storm event to either the hour at which the runoff has receded to the same value that existed before the storm, or alternately to the hour immediately preceding the next storm event.

The number of storms selected for the Terra Alta, Elkins, Beckley, and Dunlow gages, using the above algorithm, was 163, 137, 133, and 45, respectively. The length of the selection period for the Terra Alta, Elkins, and Beckley gages was 25 years, while Dunlow was 20 years. The distributions of the selected storms at each gage were summarized using a relative frequency histogram. Each hourly accumulation within each storm was normalized using the total depth for that storm. These normalized hourly values were then averaged over the total number of storms and plotted on the same axes, providing a combined measure of storm distribution by relative depth that occurs in each hour of the storm for all selected storms at each gage, as shown in Figure 11. Since most of the selected storm events were of relatively short duration, the distributions in Figure 11 are heavily weighted in the first few hours of the 24 hour storm window.

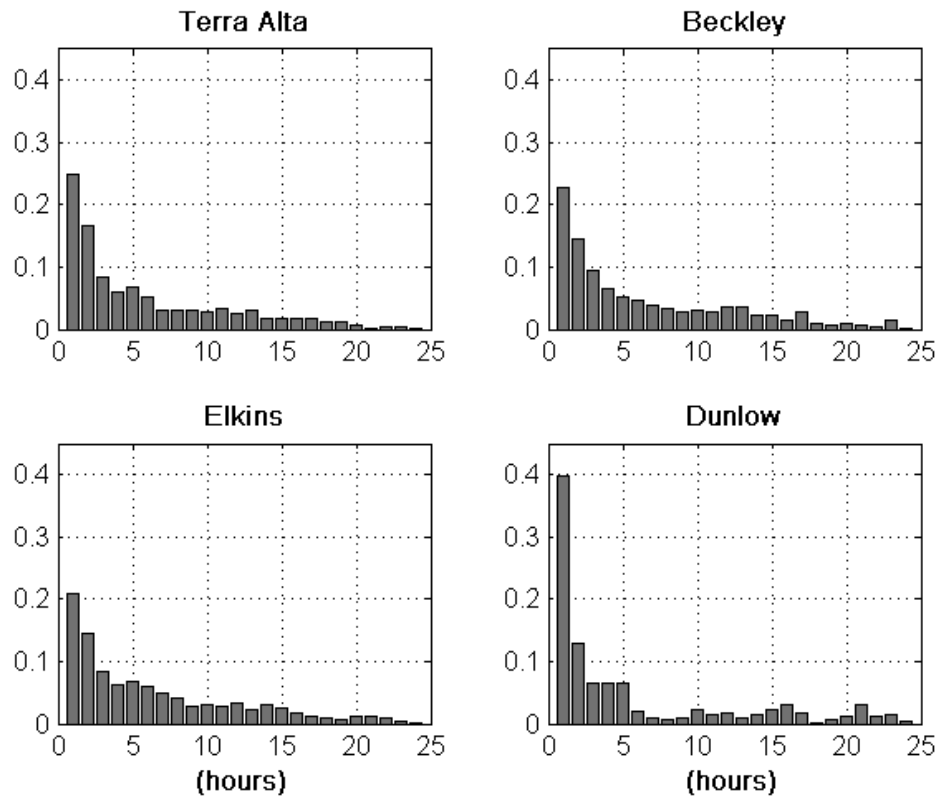


Figure 11. Storm Distribution Relative Frequency Histograms for each Gage

3.0 Results and Discussion

3.1 Cyclic Method

Figures 12-27 are plots of the computed CN's vs. Storm Depth for Clay Loam and Silt Loam soils, using initial abstraction ratio (λ) values of 0.05 and 0.2 over the full range of soil and storm depths for each storm distribution (see Table 3). Figures 12-15 are plots of the Clay Loam soil with $\lambda = 0.2$. Figures 16-19 are plots of the Silt Loam soil with $\lambda = 0.2$. Figures 20-23 are plots of the Clay Loam soil with $\lambda = 0.05$, and Figures 24-27 are plots of the Silt Loam soil with $\lambda = 0.05$.

By examining Figures 12-27, it is apparent that the CN is a function of soil type, soil depth, storm depth, and storm distribution. In all cases, the Clay Loam soil results in higher CN values than the Silt Loam, due to the lower infiltration capacity of the Clay Loam as governed by the INFILT parameter (Table 2). For each soil type and storm distribution, CN values generally decrease with increasing soil depth. This can be explained by the increase in the upper and lower zone soil moisture storage parameters (UZSN and LZSN) with the increase in soil depth (Equations 10 and 12). For each soil depth, the CN's also vary with storm depth, typically describing violent behavior where the CN's increase abruptly and approach a constant value as storm depth increases. As Hawkins (1993) noted, the violent behavior is often preceded by complacent behavior at lower storm depths as shown in Figures 12-14, 18, 20-22, and 24-26. Additionally, the violent behavior is more prevalent at lesser soil depths and trends toward standard behavior as soil depth increases (Figures 12-14, 18, and 20-27).

The CN values are also dependent on the storm distributions. The curves of the CN's vs. storm depth for the Uniform and Full Triangular distributions describe a smoother shape than those from the Type II and WDM Triangular distributions. This may be explained by observing that the Type II and WDM Triangular distributions deliver the majority of the total storm depth in a period of time that is less than two hours (Figure 7). Conversely, the Uniform and Full Triangular distributions allocate rainfall more uniformly throughout the twenty-four hour period, resulting in lower hourly intensities.

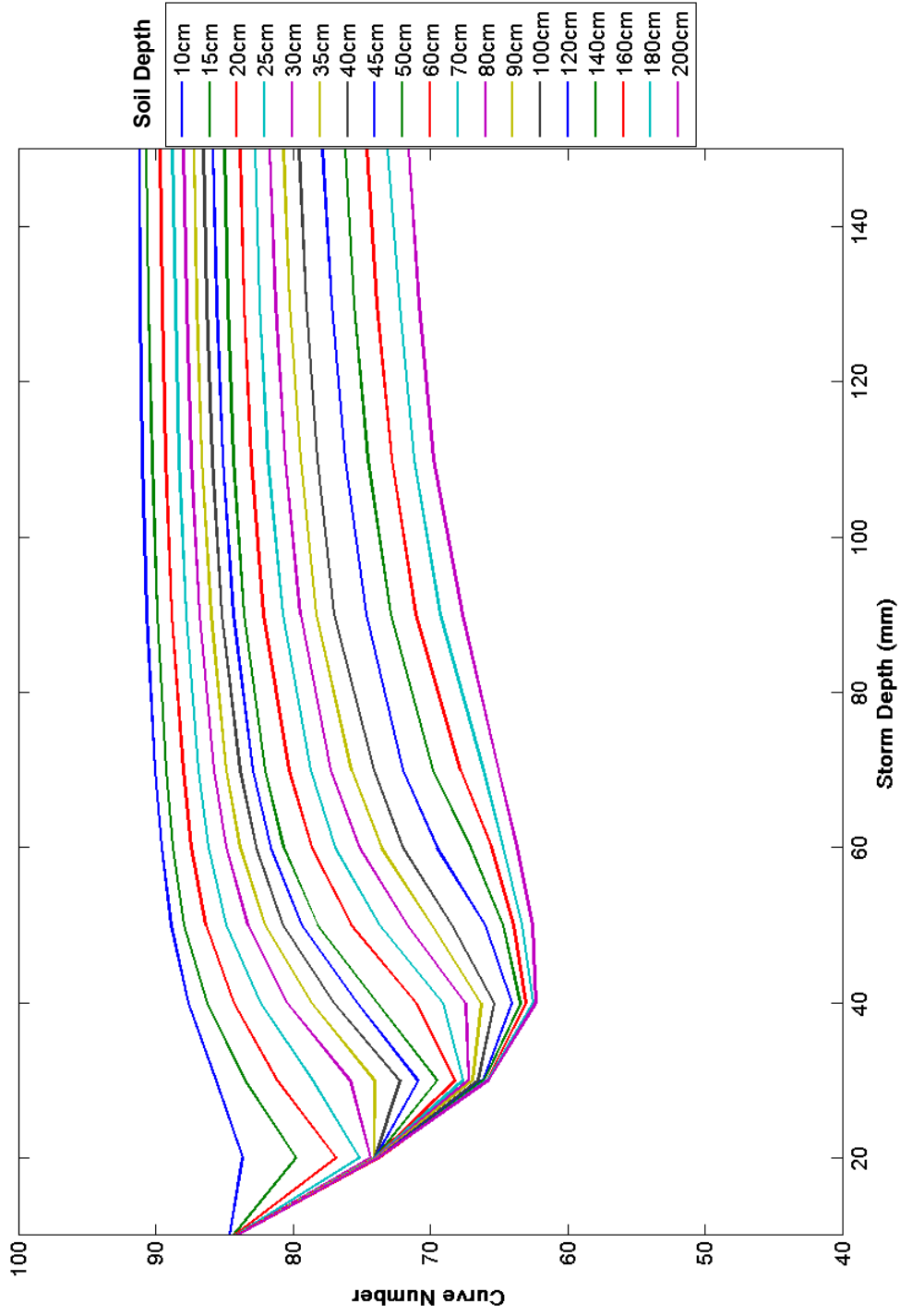


Figure 12. Cyclic Method, CN vs. Storm Depth, Clay Loam, $\lambda = 0.2$, Uniform Storm Distribution

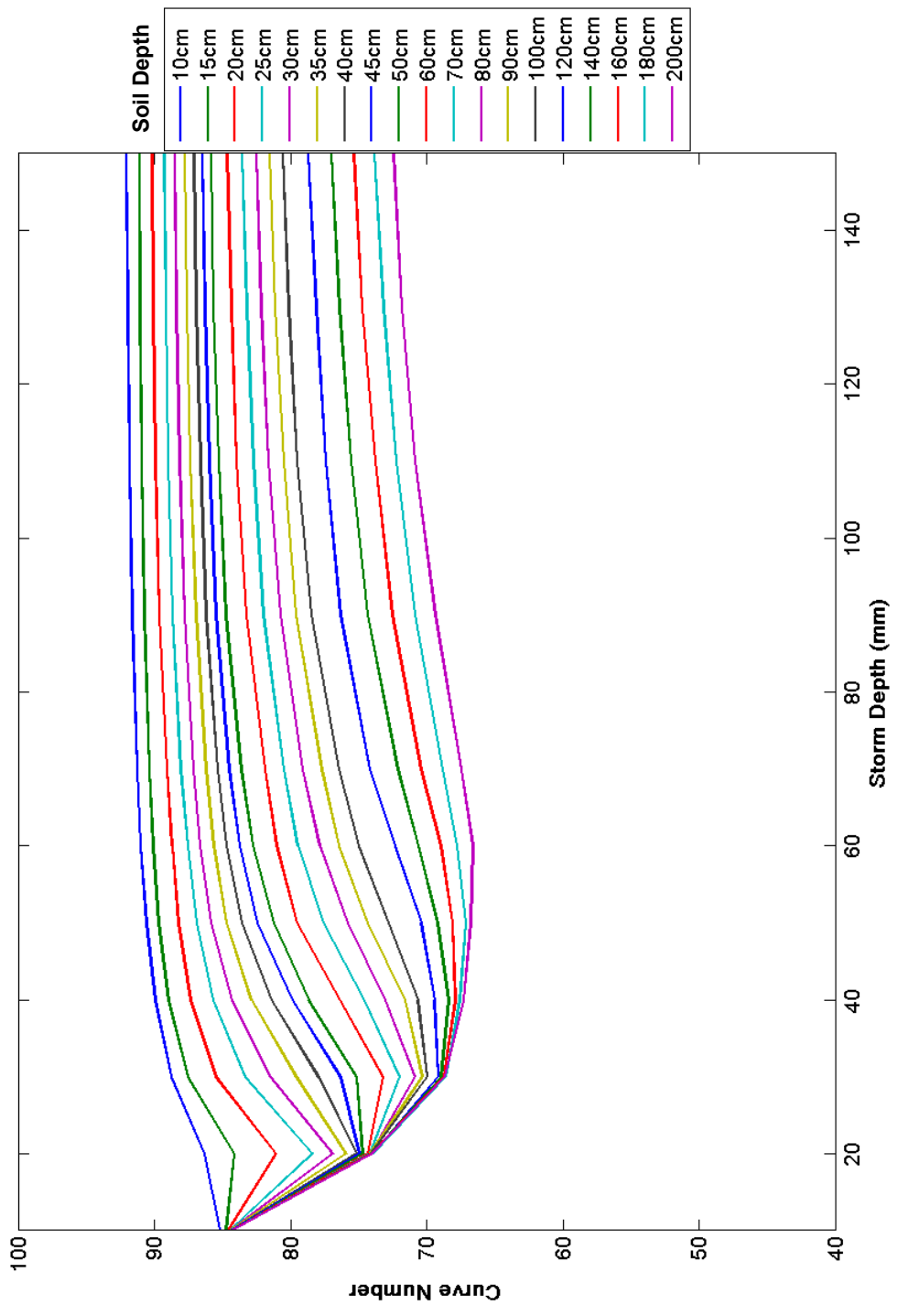


Figure 13. Cyclic Method, CN vs. Storm Depth, Clay Loam, $\lambda = 0.2$, Full Triangular Storm Distribution

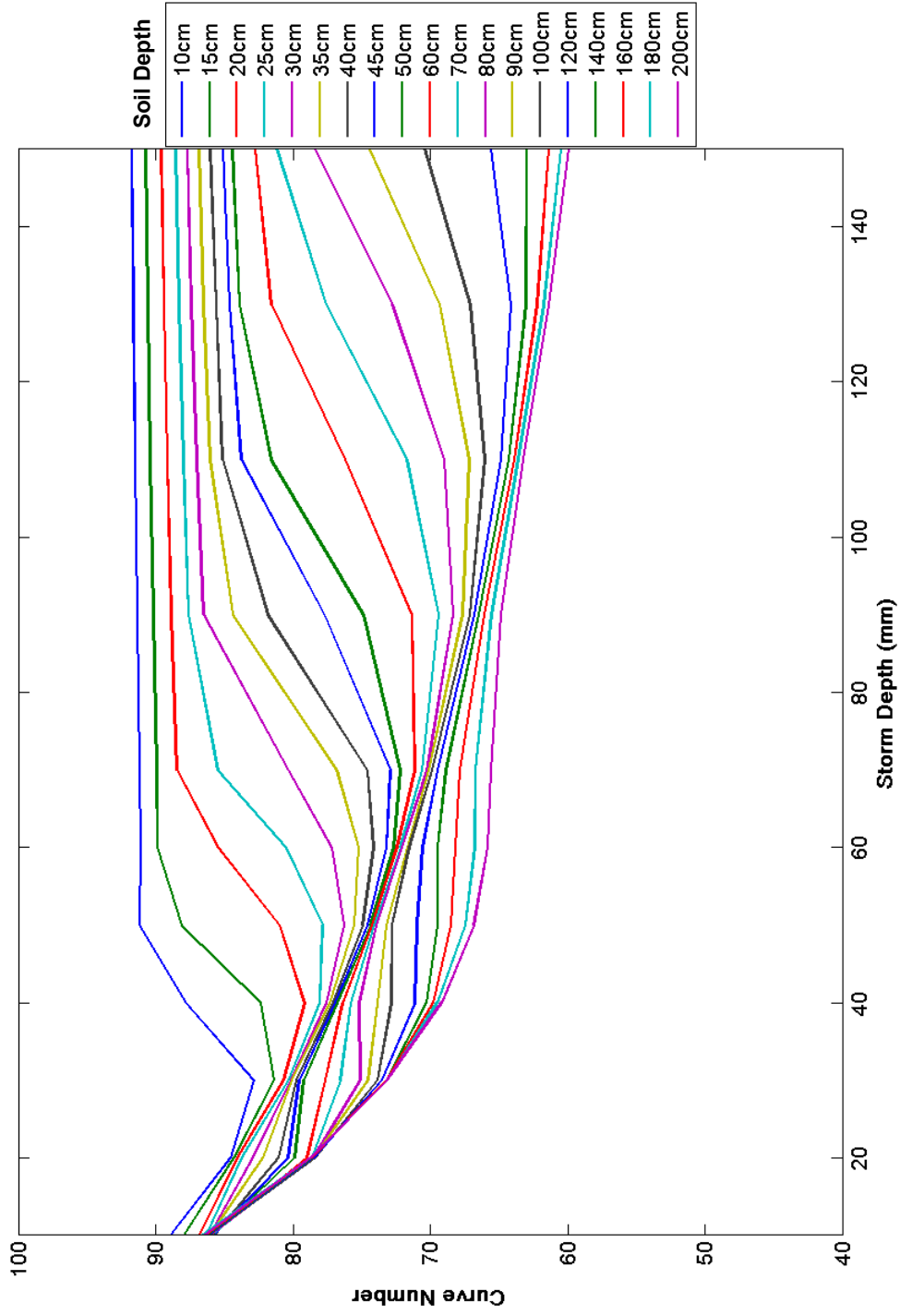


Figure 14. Cyclic Method, CN vs. Storm Depth, Clay Loam, $\lambda = 0.2$, Type II Storm Distribution

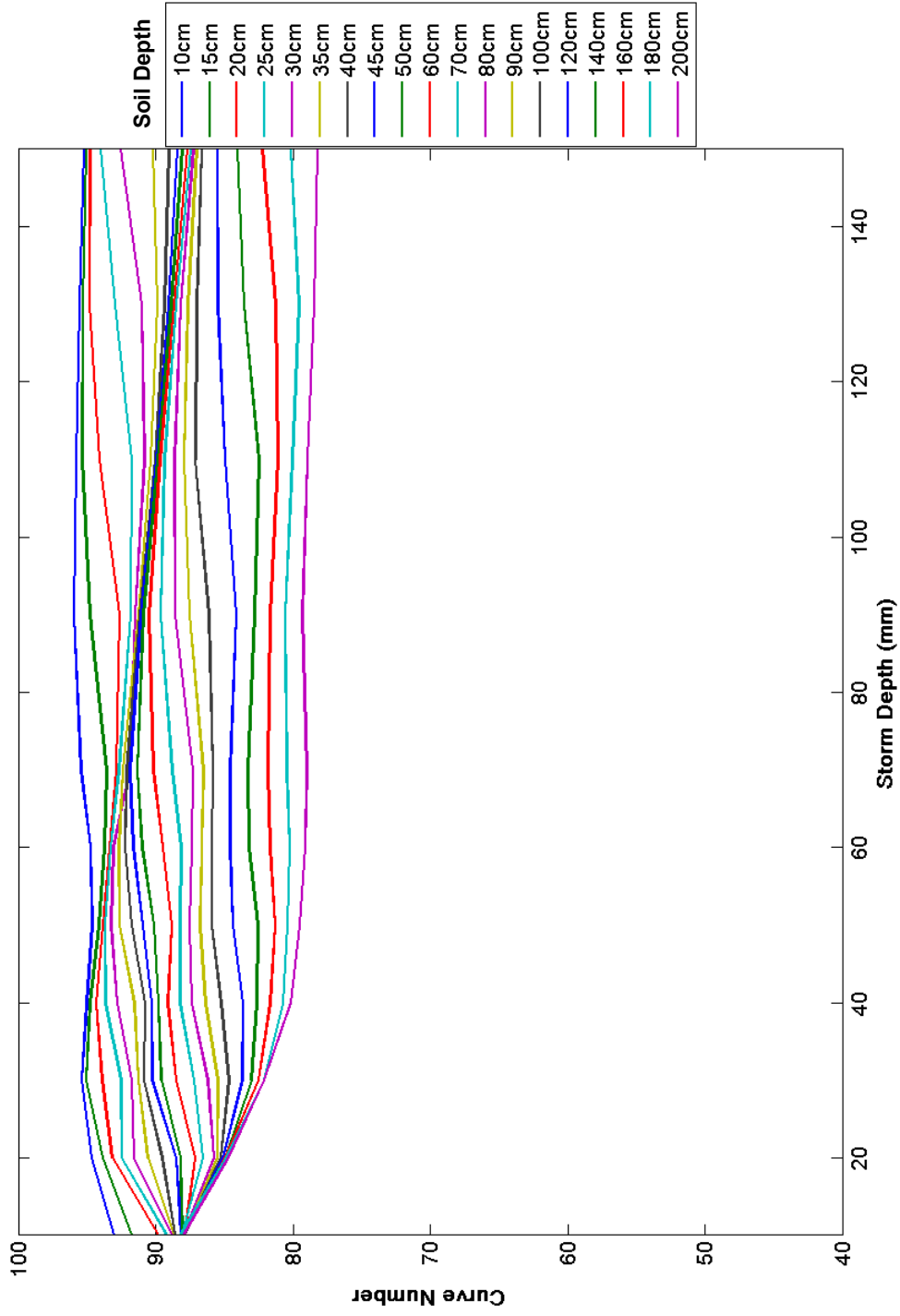


Figure 15. Cyclic Method, CN vs. Storm Depth, Clay Loam, $\lambda = 0.2$, WDM Triangular Storm Distribution

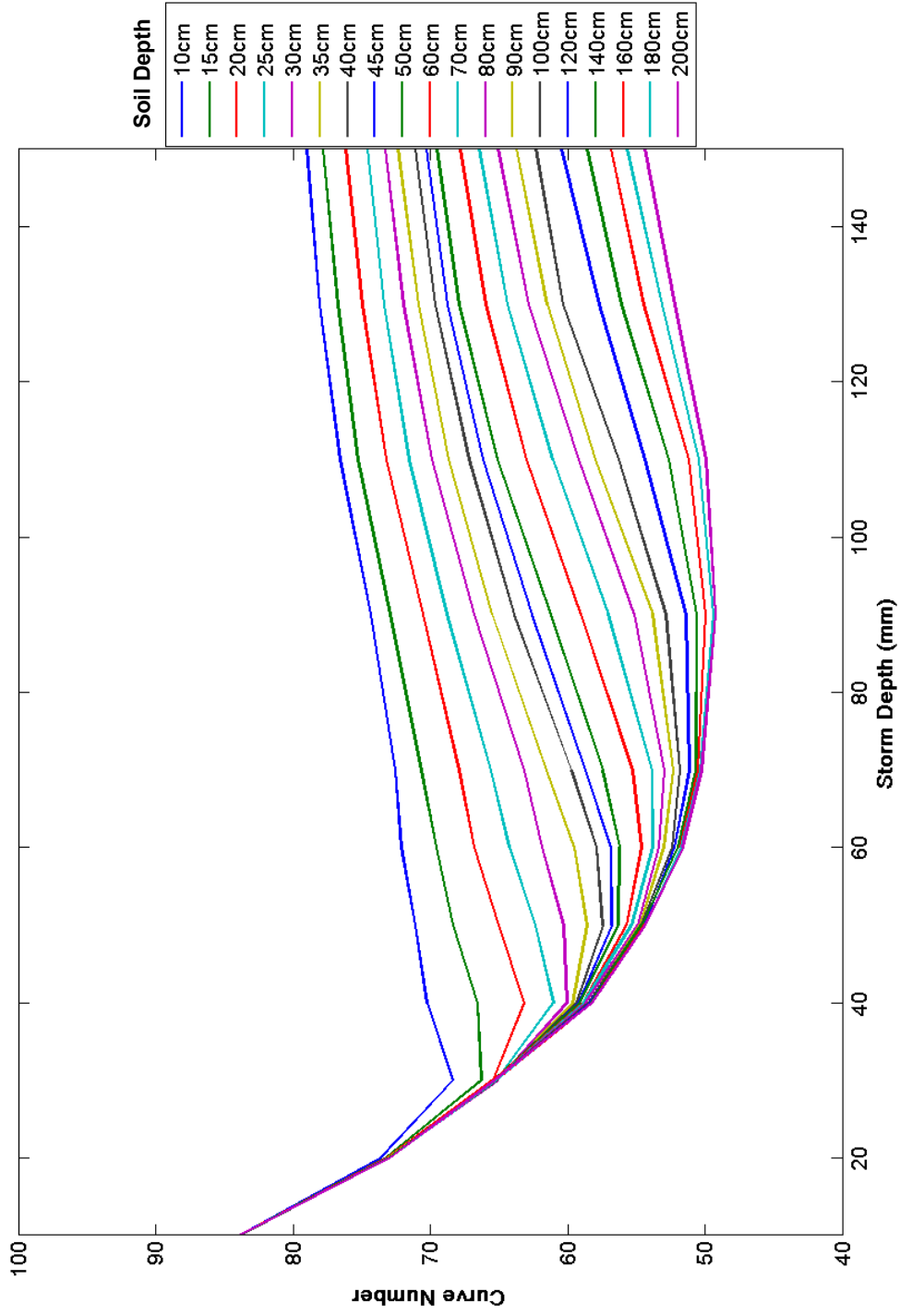


Figure 16. Cyclic Method, CN vs. Storm Depth, Silt Loam, $\lambda = 0.2$, Uniform Storm Distribution

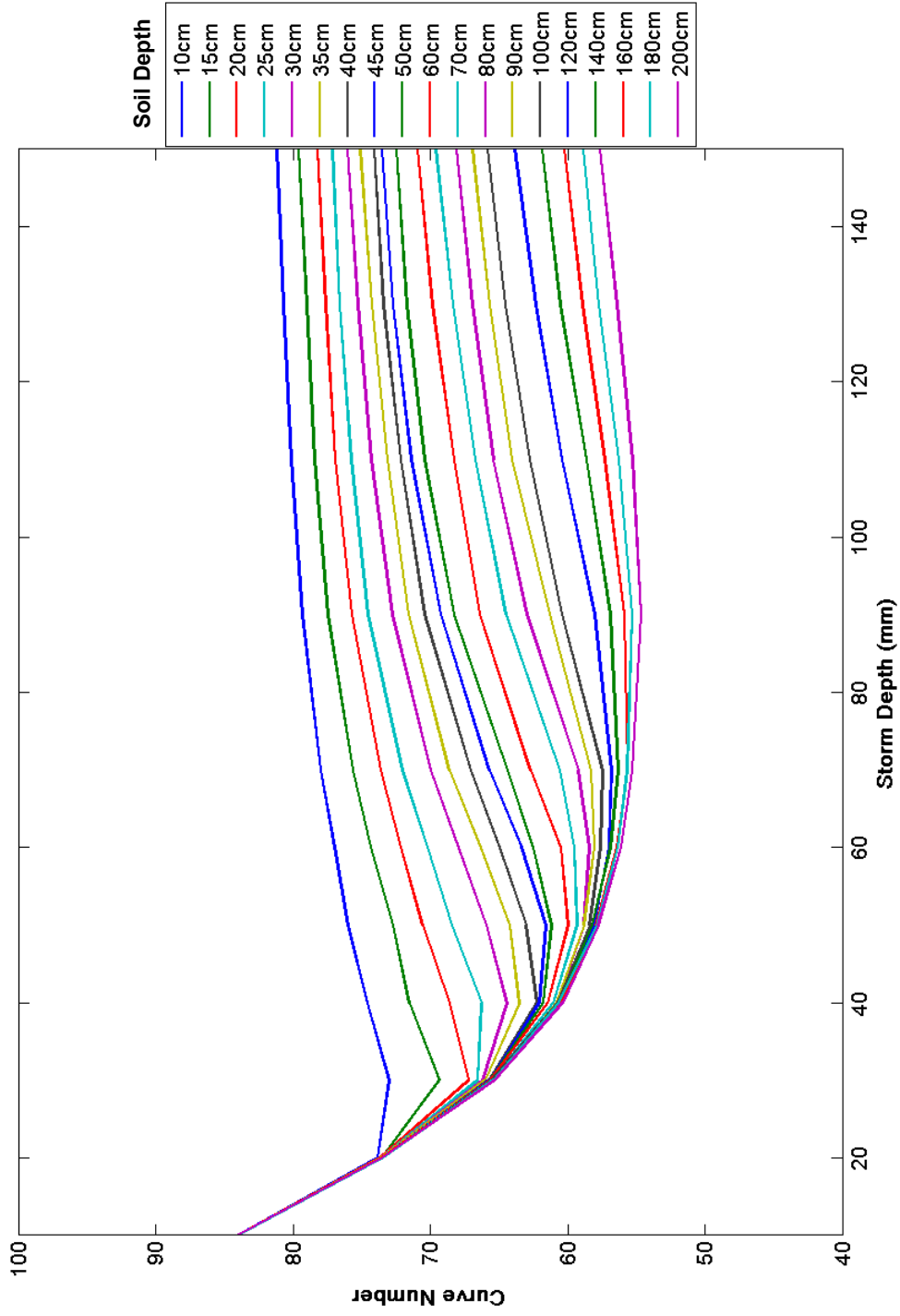


Figure 17. Cyclic Method, CN vs. Storm Depth, Silt Loam, $\lambda = 0.2$, Full Triangular Storm Distribution

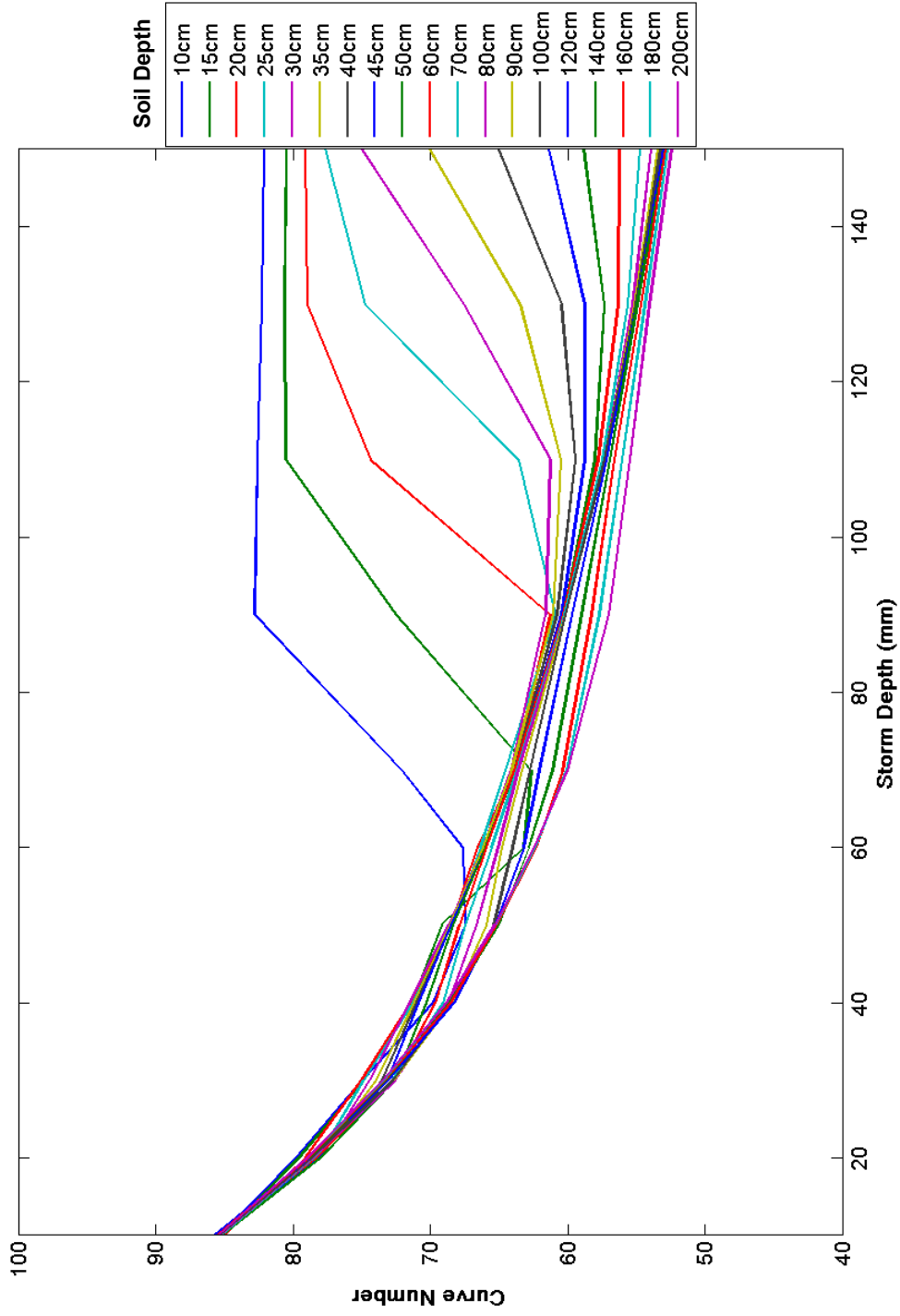


Figure 18. Cyclic Method, CN vs. Storm Depth, Silt Loam, $\lambda = 0.2$, Type II Storm Distribution

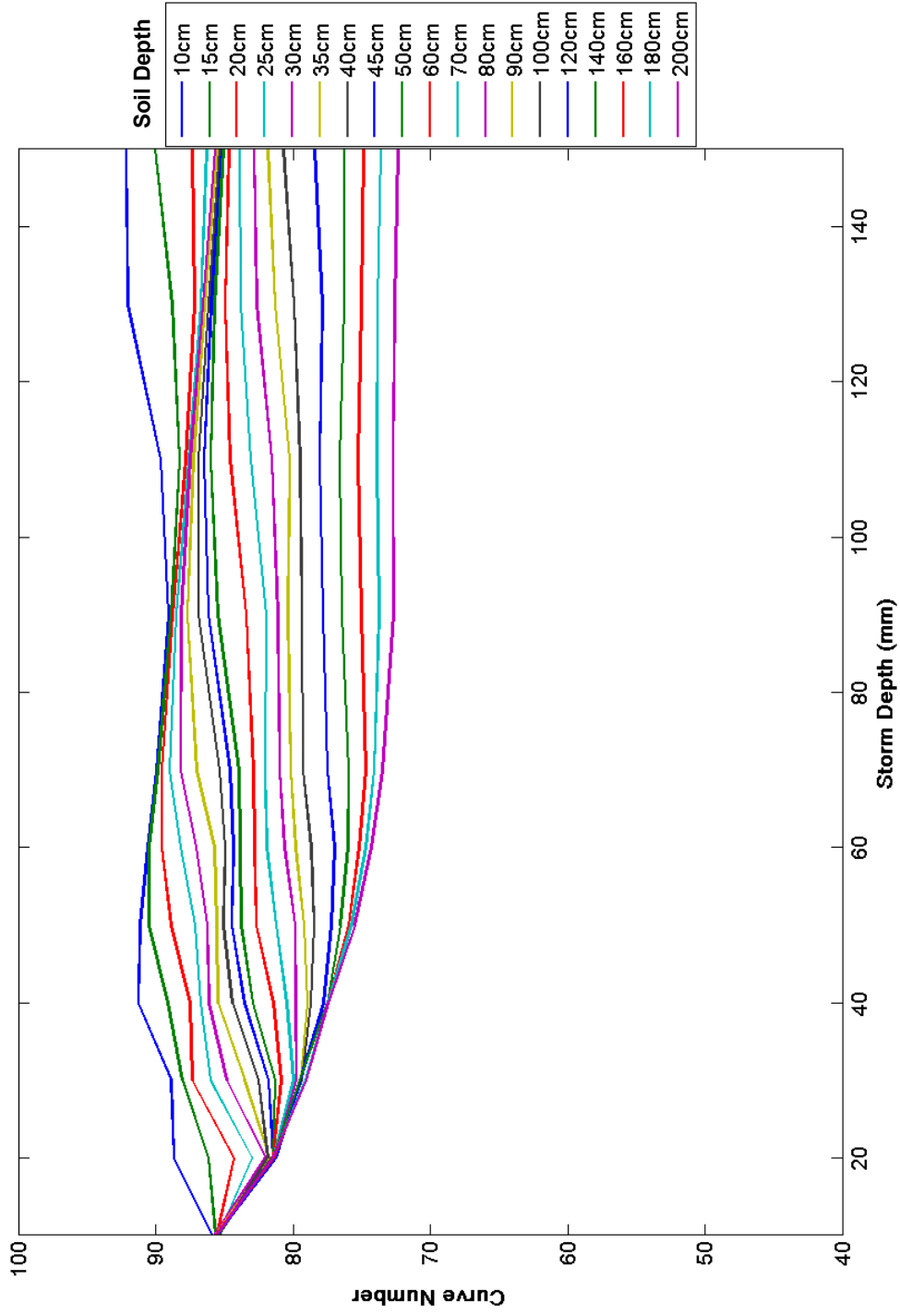


Figure 19. Cyclic Method, CN vs. Storm Depth, Silt Loam, $\lambda = 0.2$, WDM Triangular Storm Distribution

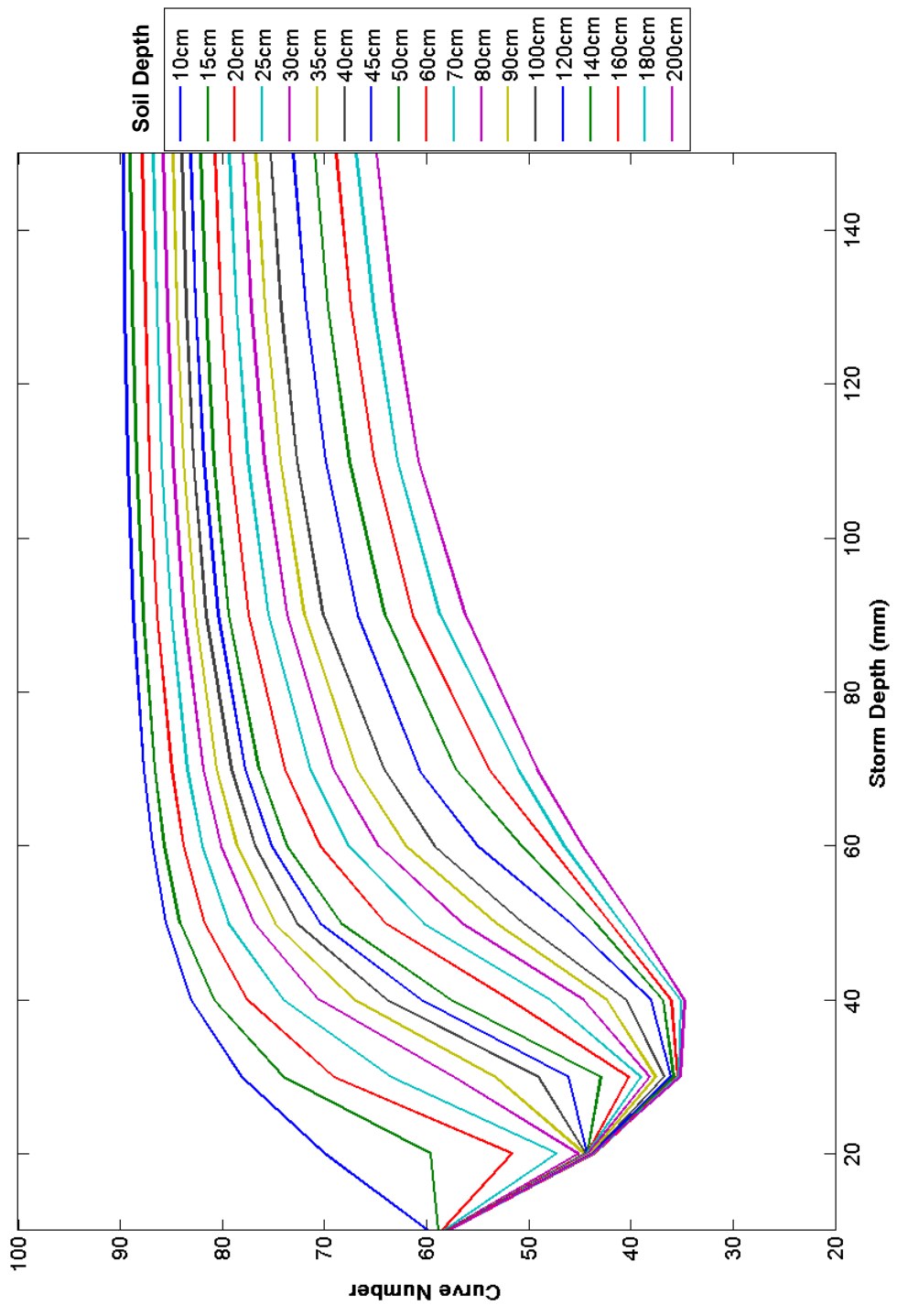


Figure 20. Cyclic Method, CN vs. Storm Depth, Clay Loam, $\lambda = 0.05$, Uniform Storm Distribution

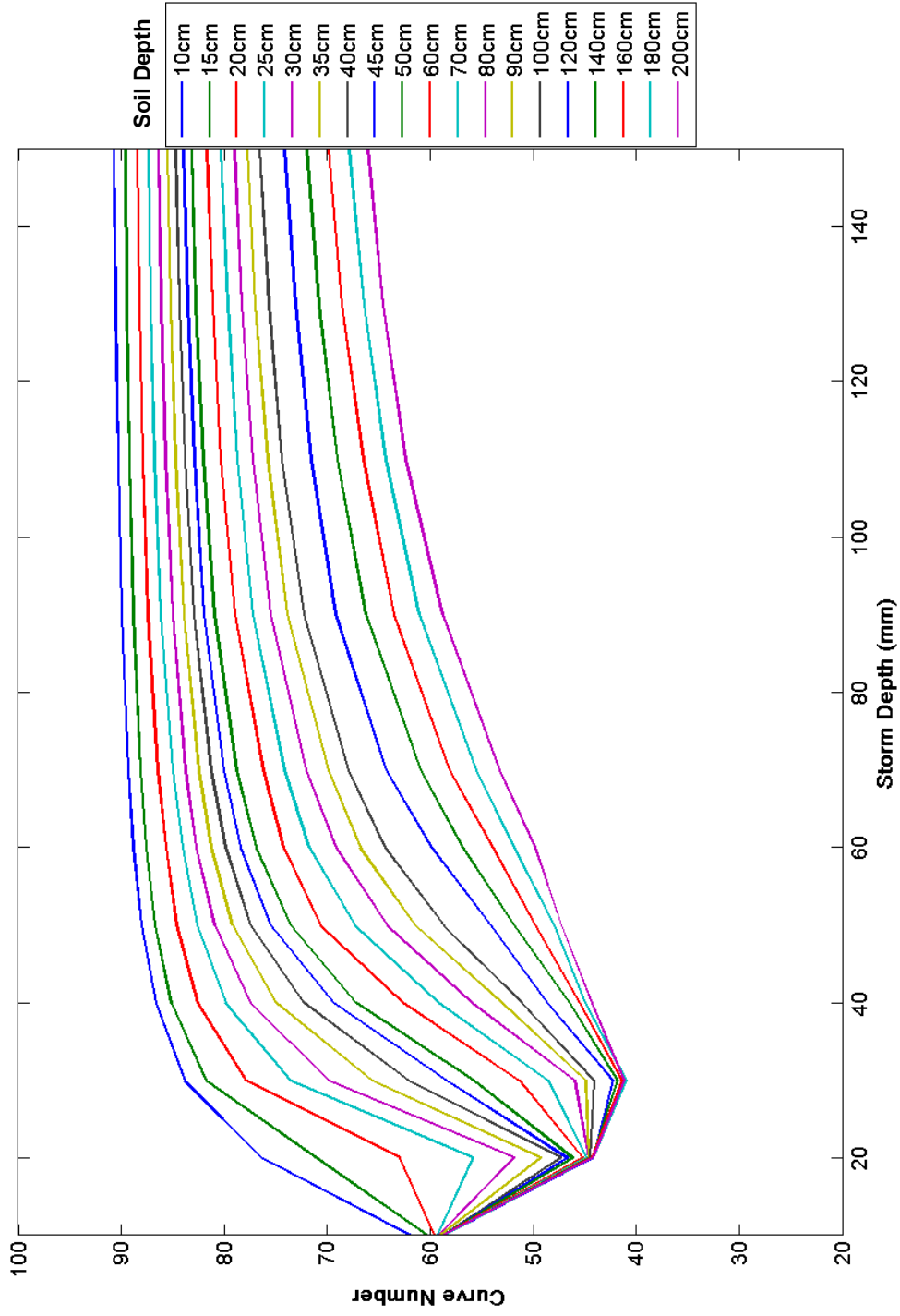


Figure 21 . Cyclic Method, CN vs. Storm Depth, Clay Loam, $\lambda = 0.05$, Full Triangular Storm Distribution

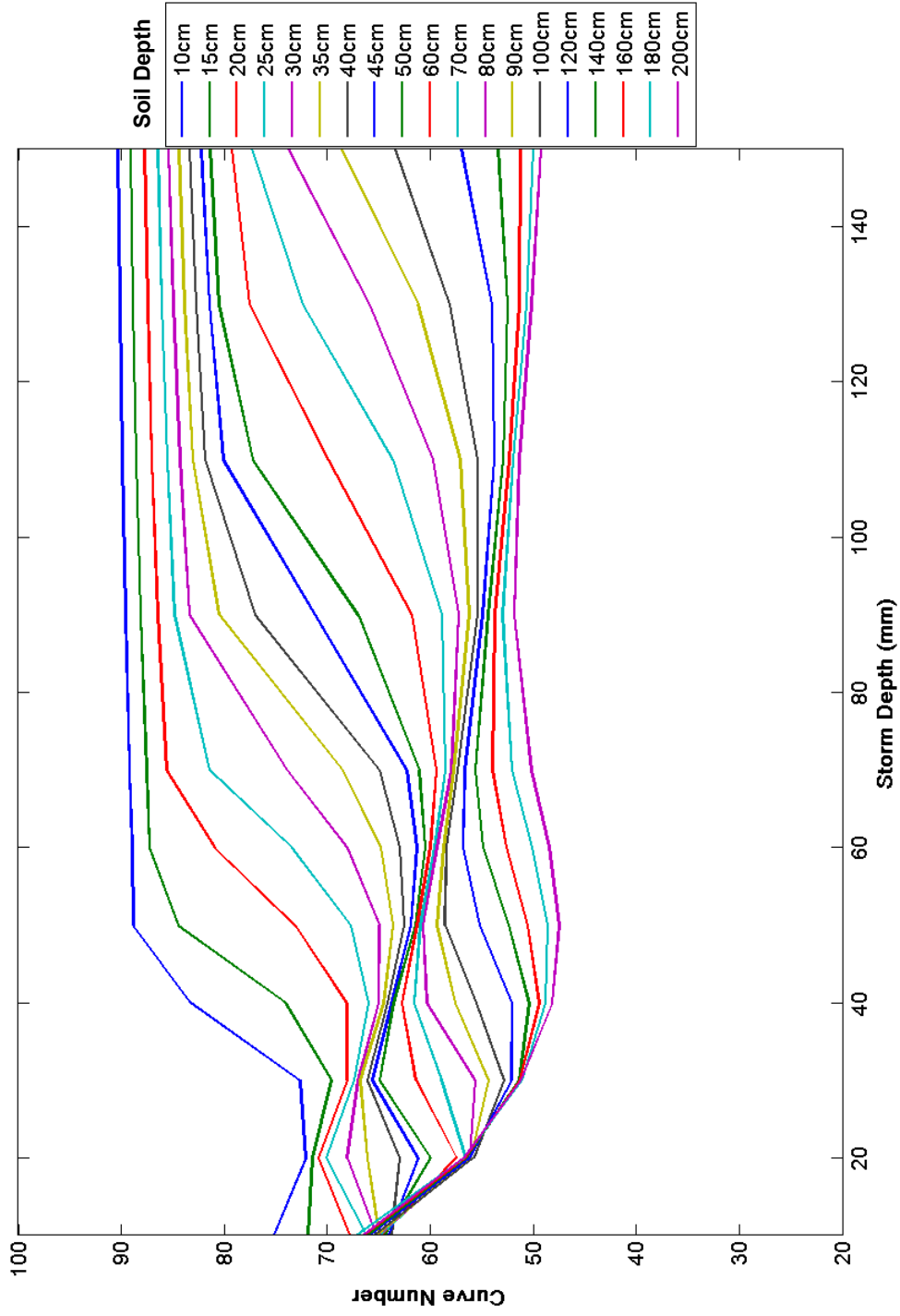


Figure 22. Cyclic Method, CN vs. Storm Depth, Clay Loam, $\lambda = 0.05$, Type II Storm Distribution

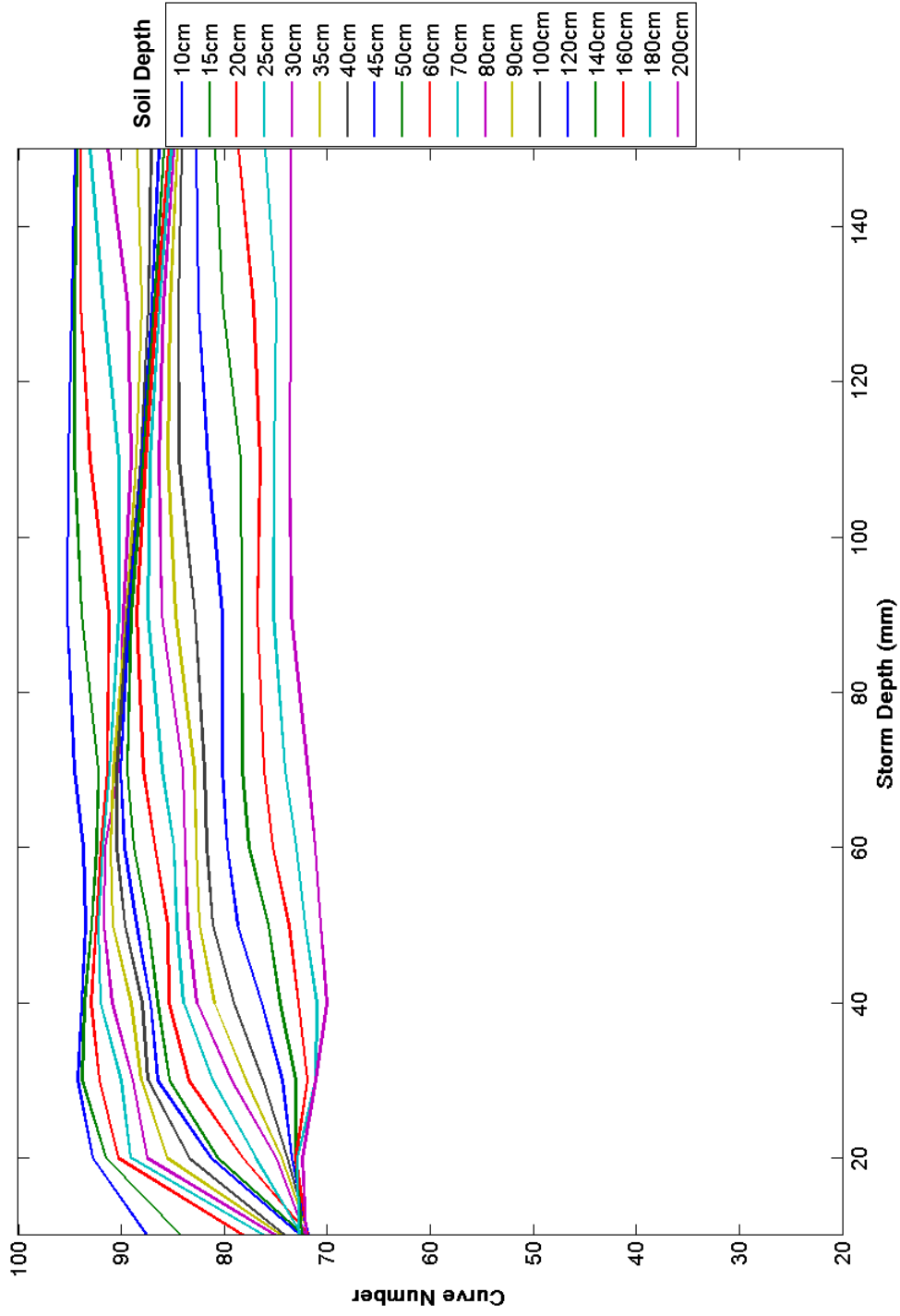


Figure 23. Cyclic Method, CN vs. Storm Depth, Clay Loam, $\lambda = 0.05$, WDM Triangular Storm Distribution

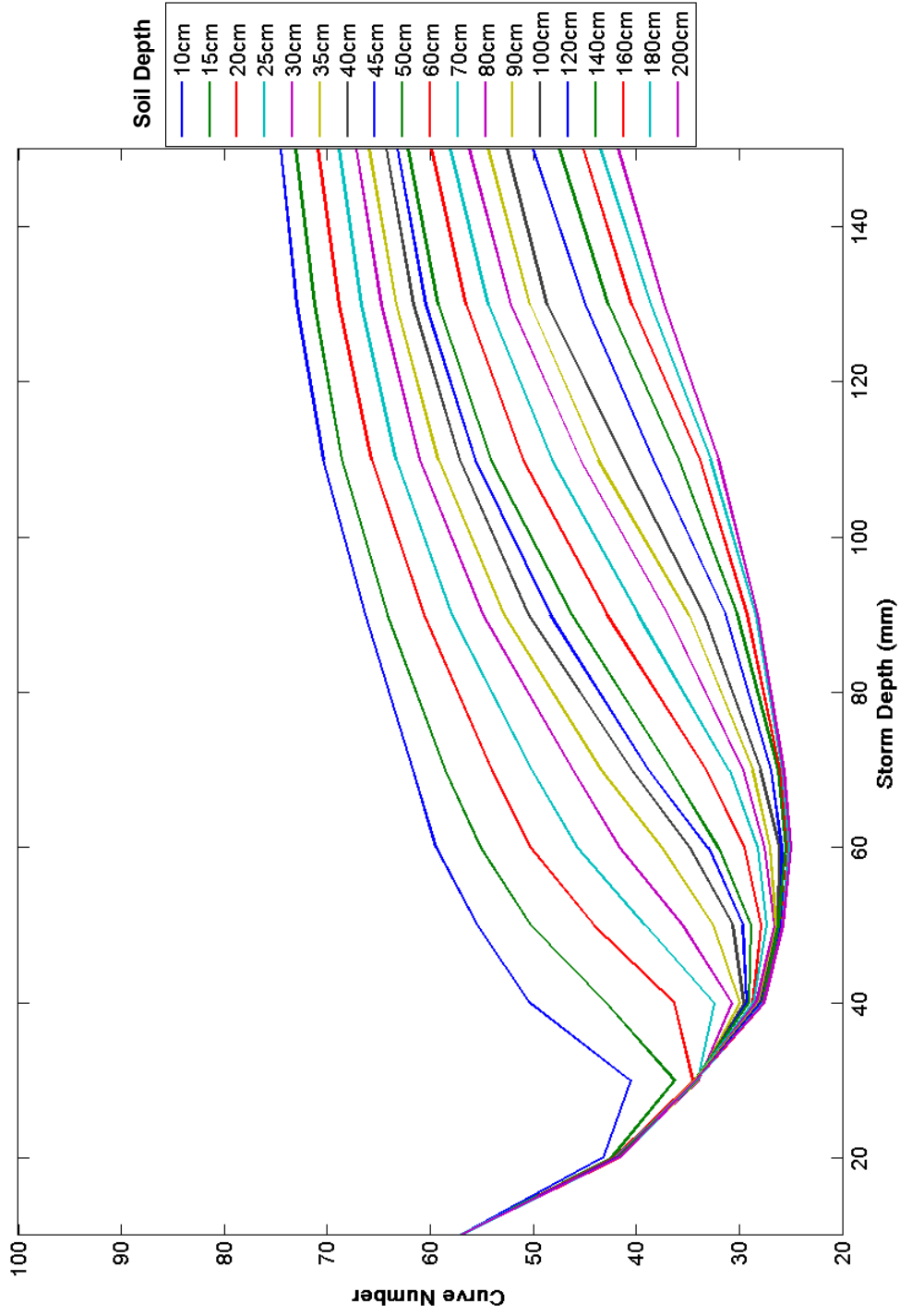


Figure 24. Cyclic Method, CN vs. Storm Depth, Silt Loam, $\lambda = 0.05$, Uniform Storm Distribution

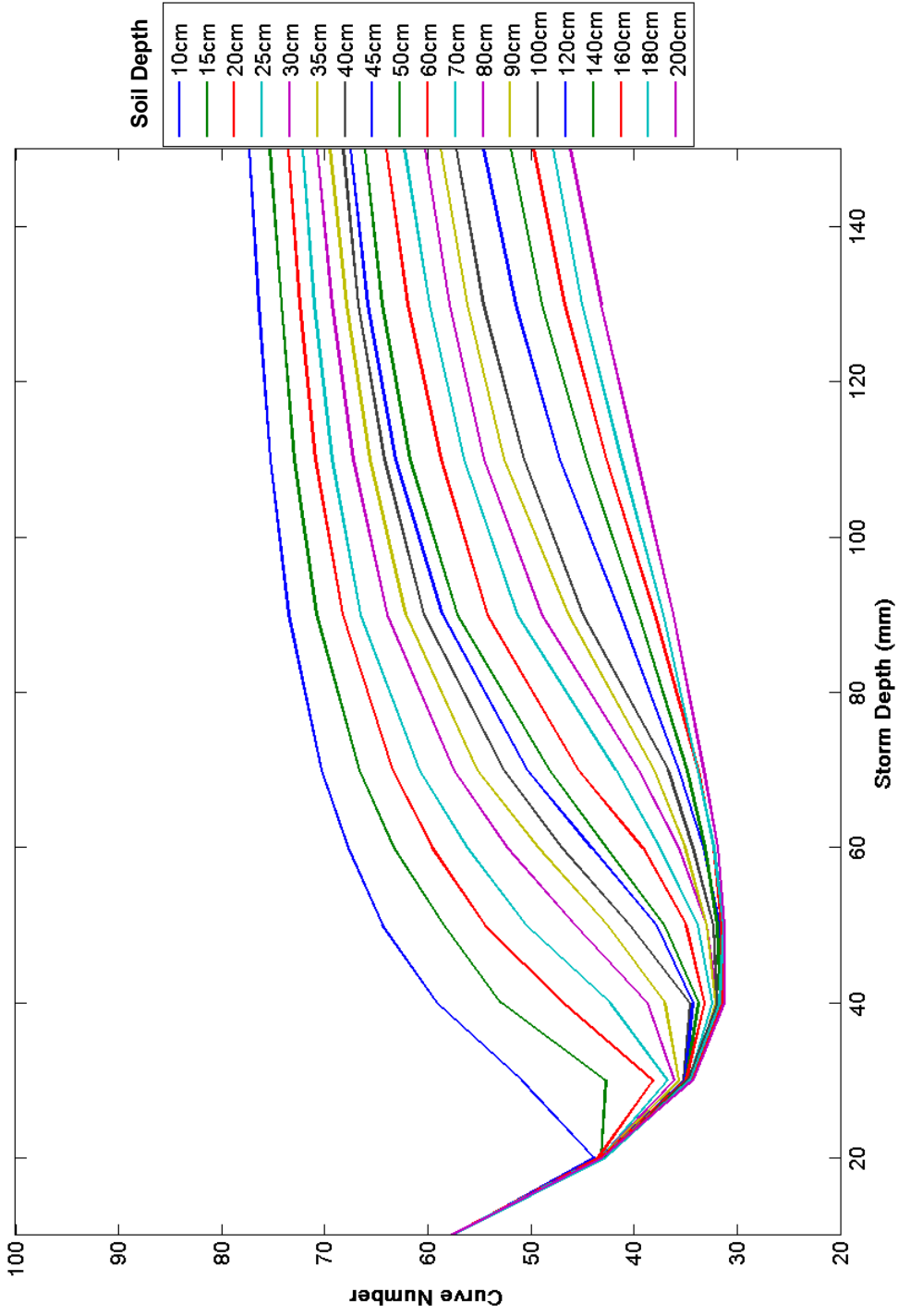


Figure 25. Cyclic Method, CN vs. Storm Depth, Silt Loam, $\lambda = 0.05$, Full Triangular Storm Distribution

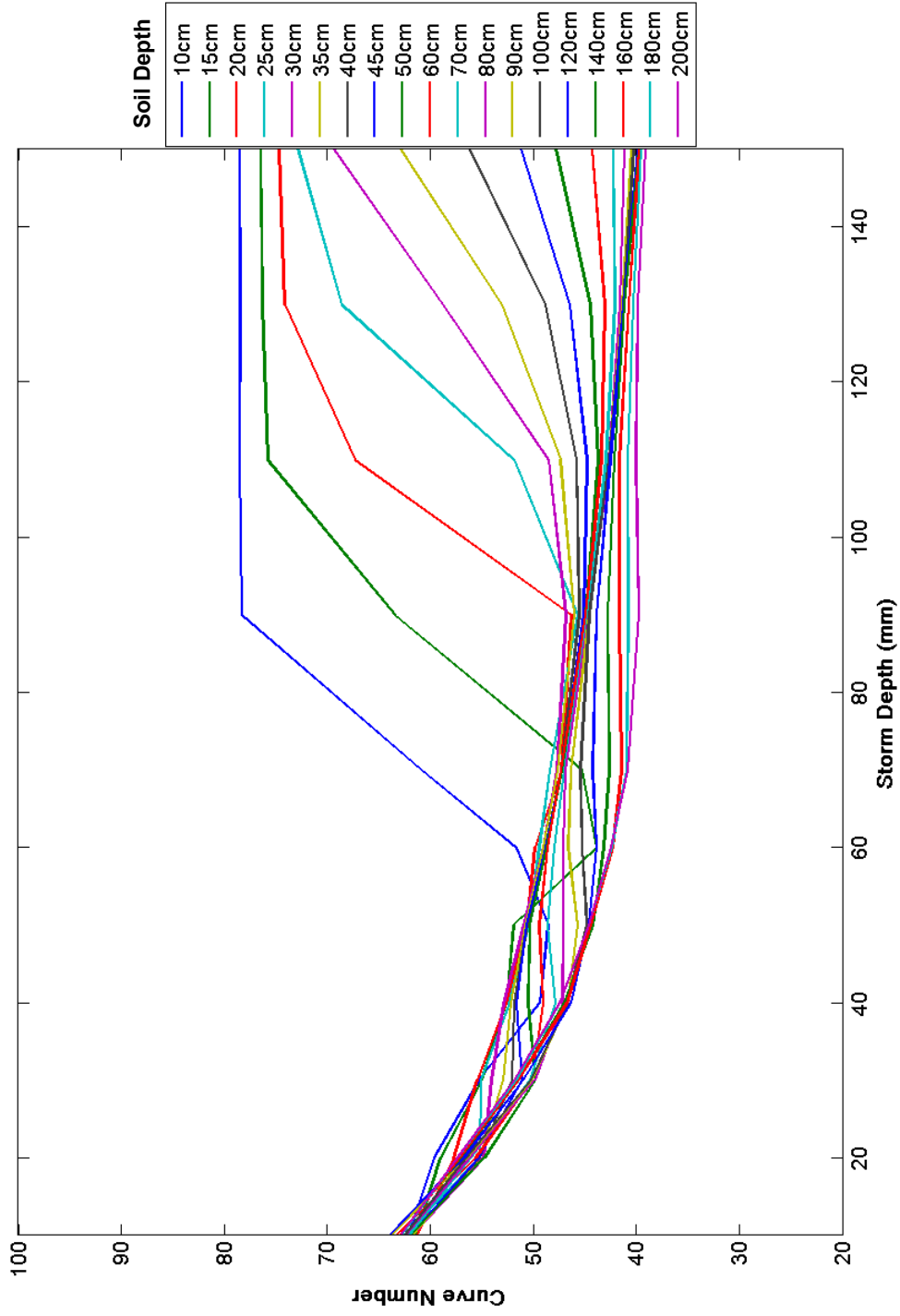


Figure 26. Cyclic Method, CN vs. Storm Depth, Silt Loam, $\lambda = 0.05$, Type II Storm Distribution

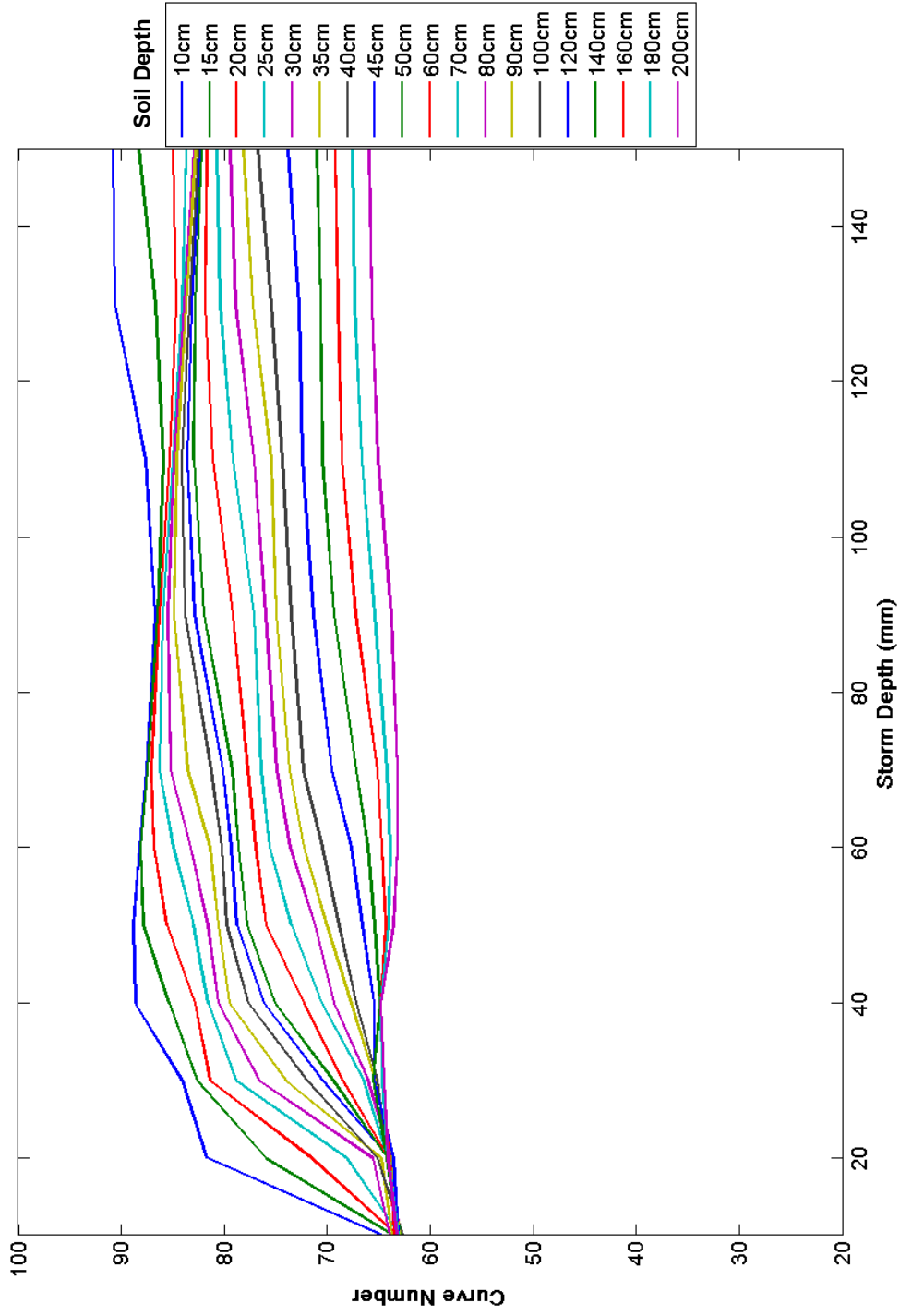


Figure 27. Cyclic Method, CN vs. Storm Depth, Silt Loam, $\lambda = 0.05$, WDM Triangular Storm Distribution

The WDM Triangular distribution results in the highest overall CN values, as shown in Figures 15, 19, 23, and 27. This can be explained by the fact that the total twenty-four hour rainfall depth in this distribution falls within an eight hour period, therefore providing less opportunity for infiltration. To illustrate this effect, the HSPF infiltration component was accumulated from the beginning of the storm event to the beginning of the next storm event and averaged over the range of storm depths for each soil depth. The results for each storm distribution for the Clay Loam soil are shown in Figure 28. The WDM Triangular distribution results in the least amount of infiltration, therefore producing the highest CN values.

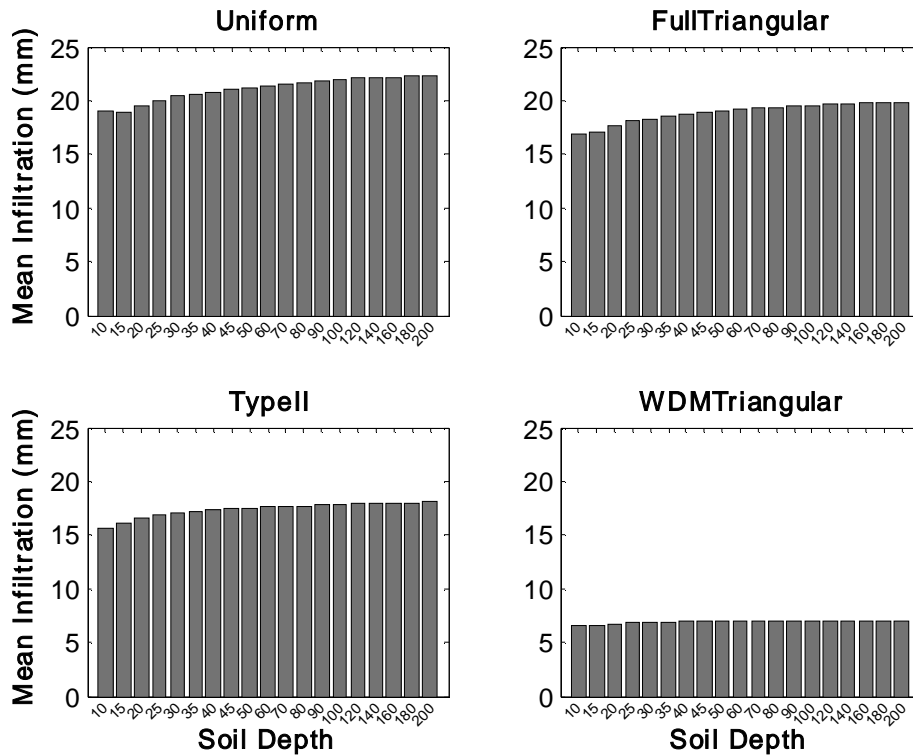


Figure 28. Cyclic Method Mean Infiltration vs. Soil Depth, Clay Loam

Finally, by comparing Figures 12-19 to Figures 20-27, it is apparent that the initial abstraction ratio (λ) value of 0.05 versus 0.2 reduces the CN's calculated at low storm depths. This makes sense physically; since the storm and runoff depths are the same for each soil depth, and less initial abstraction (corresponding to a lower λ value) results in a greater loss to the soil, producing a lower CN value. As the storm depth

increases, the initial abstraction becomes a smaller percentage of the rainfall and its effect on the CN diminishes.

The results of the Cyclic Method indicate that the storm distributions consisting of high intensity hourly events (Type II and WDM Triangular) produce more irregular variation of the CN with storm depth (for example, see Figures 12-15) . The curves of the lesser soil depths tend to describe a violent behavior, trending toward standard behavior as soil depth increases (Figures 12-25). Finally, the WDM Triangular distribution resulted in the lowest infiltration depth and therefore the highest CN values (Figures 15, 19, 23, 27, and 28). These findings can be compared to those of the Asymptotic Method that follows.

3.2 Asymptotic Method

The Asymptotic Method was applied to the same range of soil types and soil depths as the Cyclic method. The CN's calculated from the ranked rainfall-runoff pairs were plotted against rainfall for each soil depth. The equation that resulted in the highest R-squared value (violent, equation 15, or standard, equation 16) was fit to the data by minimizing the least-squared error using a Matlab (Mathworks, Inc.) optimization function (Lagarias, 1998). Figure 29 is an example fit of the equation for violent behavior for Clay Loam at a 20 cm soil depth with a simulation using data from the Terra Alta gage.

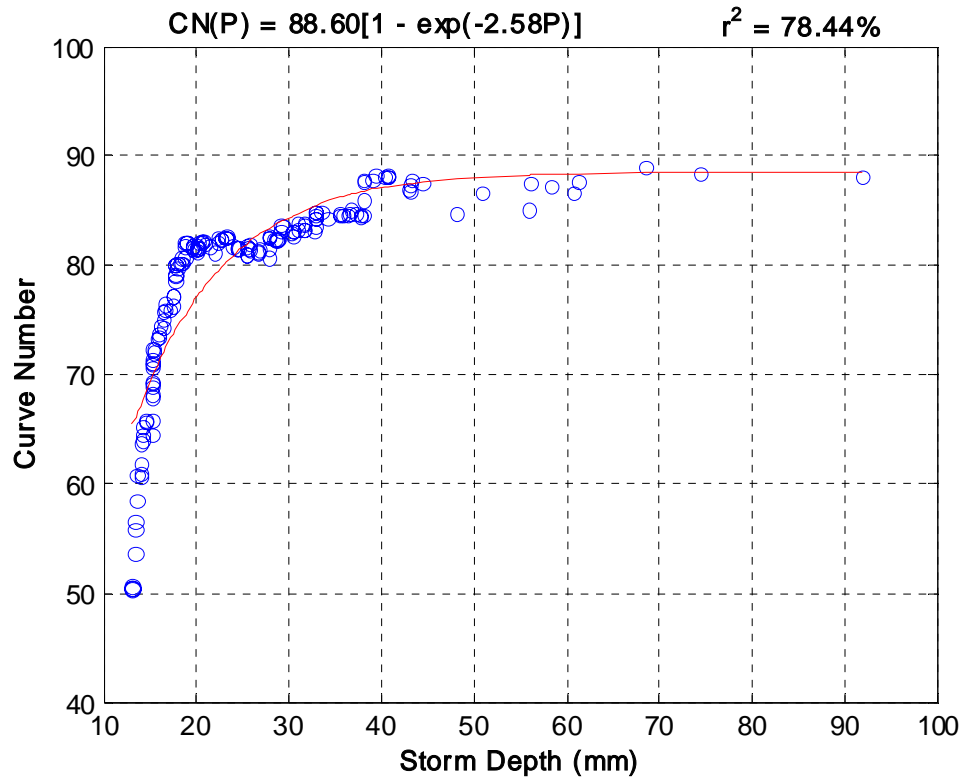


Figure 29. Violent Curve Fit, Clay Loam, 20 cm, Terra Alta gage, $\lambda = 0.05$

In this case the abrupt rise in CN, characteristic of the so-called violent behavior is obvious and equation 17 fits the data reasonably well ($R^2 = 78.44\%$).

$$CN(P) = 88.60[1 - \exp(-2.58P)] \quad (17)$$

This process was repeated for each soil depth and gaging station and the resulting curves were plotted on the same axes. Certain soil depths appeared to exhibit complacent behavior at low storm depths followed by violent behavior. In these instances the complacent behavior was ignored and only the violent points were used to fit Equation 16. Figures 30 to 45 show the results for the Clay Loam and Silt Loam soils with initial abstraction ratio values of 0.05 and 0.2. Figures 30-33 are plots of the Clay Loam soil with $\lambda = 0.2$. Figures 34-37 are plots of the Silt Loam soil with $\lambda = 0.2$. Figures 38-41 are plots of the Clay Loam soil with $\lambda = 0.05$, and Figures 42-45 are plots of the Silt Loam soil with $\lambda = 0.05$.

The fits with R^2 values less than 50% are shown as dotted lines. Tables 4 and 5 show the corresponding CN_{∞} values, fitting constants, and R^2 values for each curve fit.

The type of fit is indicated with the letters 's' or 'v' (Equations 15 or 16, respectively), and the entries with R^2 values less than 50% are shown in red.

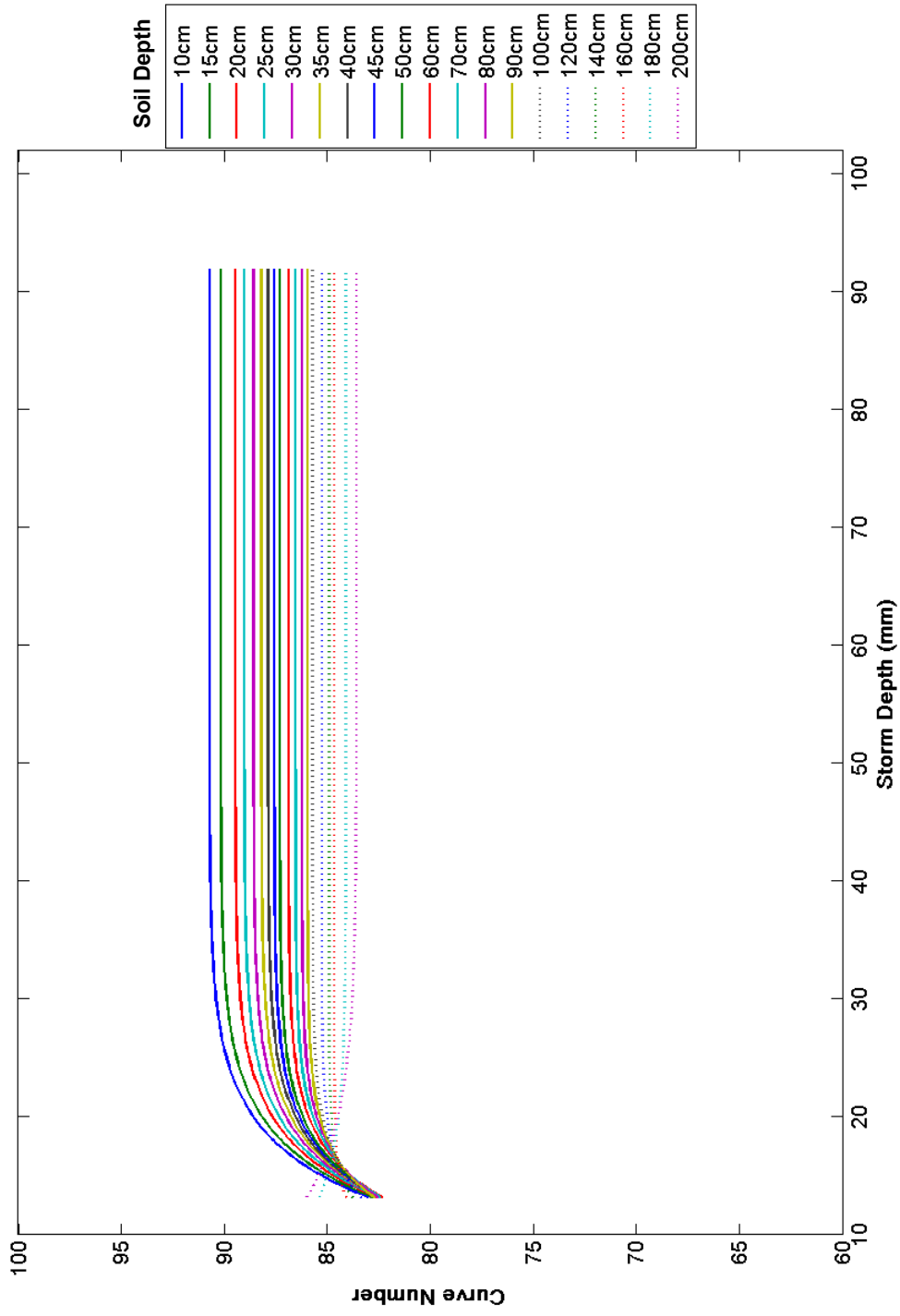


Figure 30. Asymptotic Method, CN Fit Curves vs. Storm Depth, Clay Loam, $\lambda = 0.2$, Terra Alta

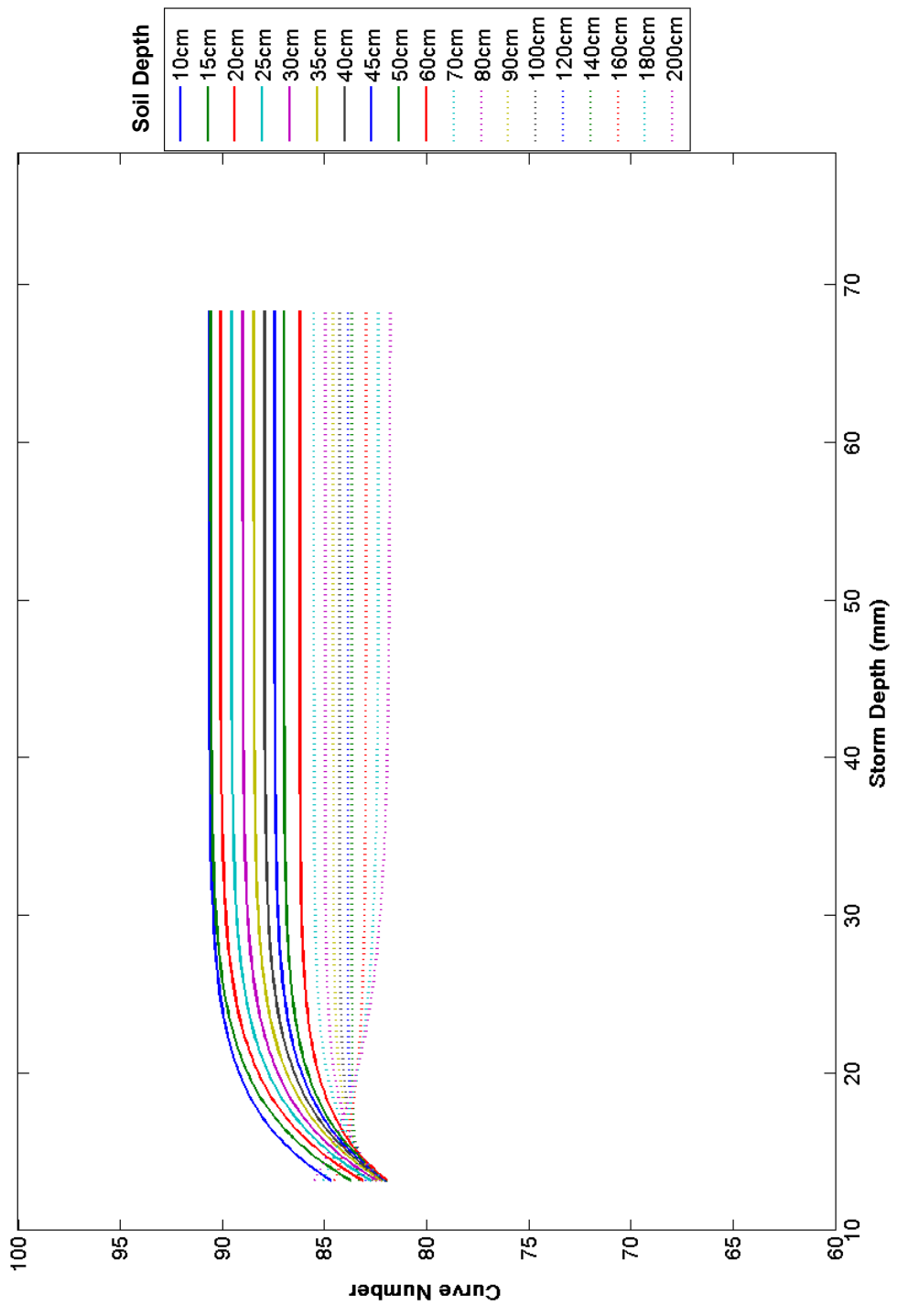


Figure 31. Asymptotic Method, CN Fit Curves vs. Storm Depth, Clay Loam, $\lambda = 0.2$, Beckley

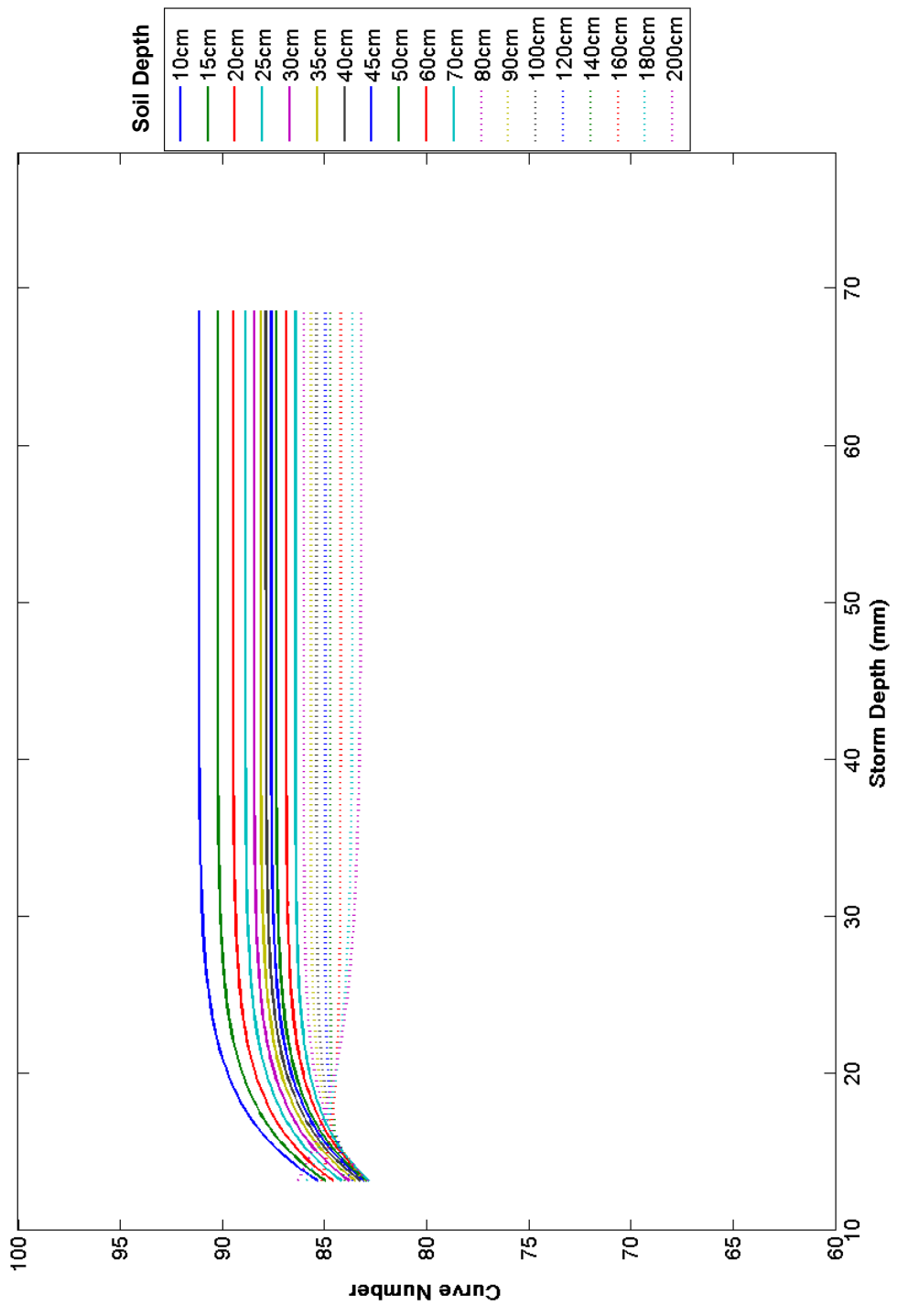


Figure 32. Asymptotic Method, CN Fit Curves vs. Storm Depth, Clay Loam, $\lambda = 0.2$, Elkins

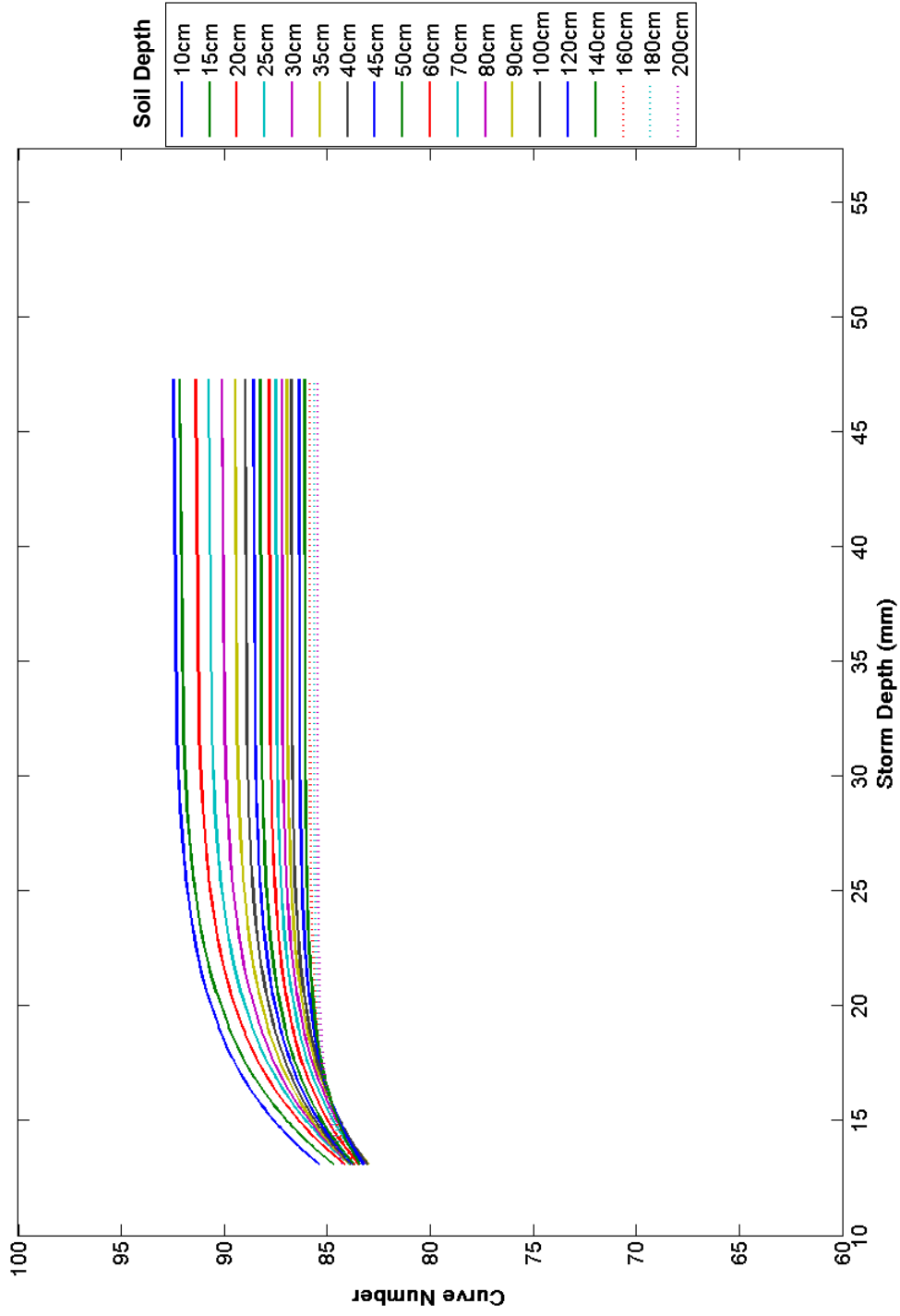


Figure 33. Asymptotic Method, CN Fit Curves vs. Storm Depth, Clay Loam, $\lambda = 0.2$, Dunlow

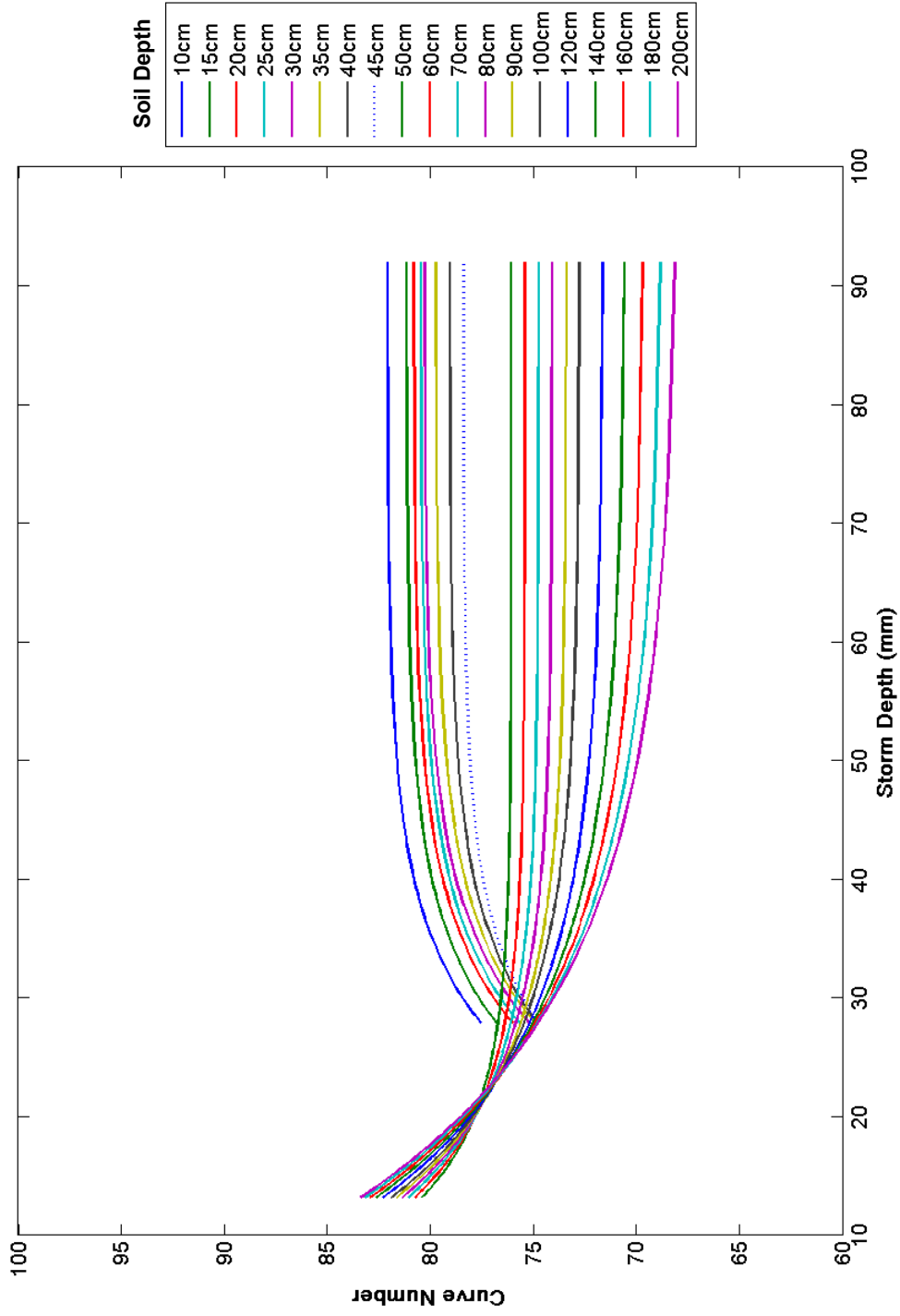


Figure 34. Asymptotic Method, CN Fit Curves vs. Storm Depth, Silt Loam, $\lambda = 0.2$, Terra Alta

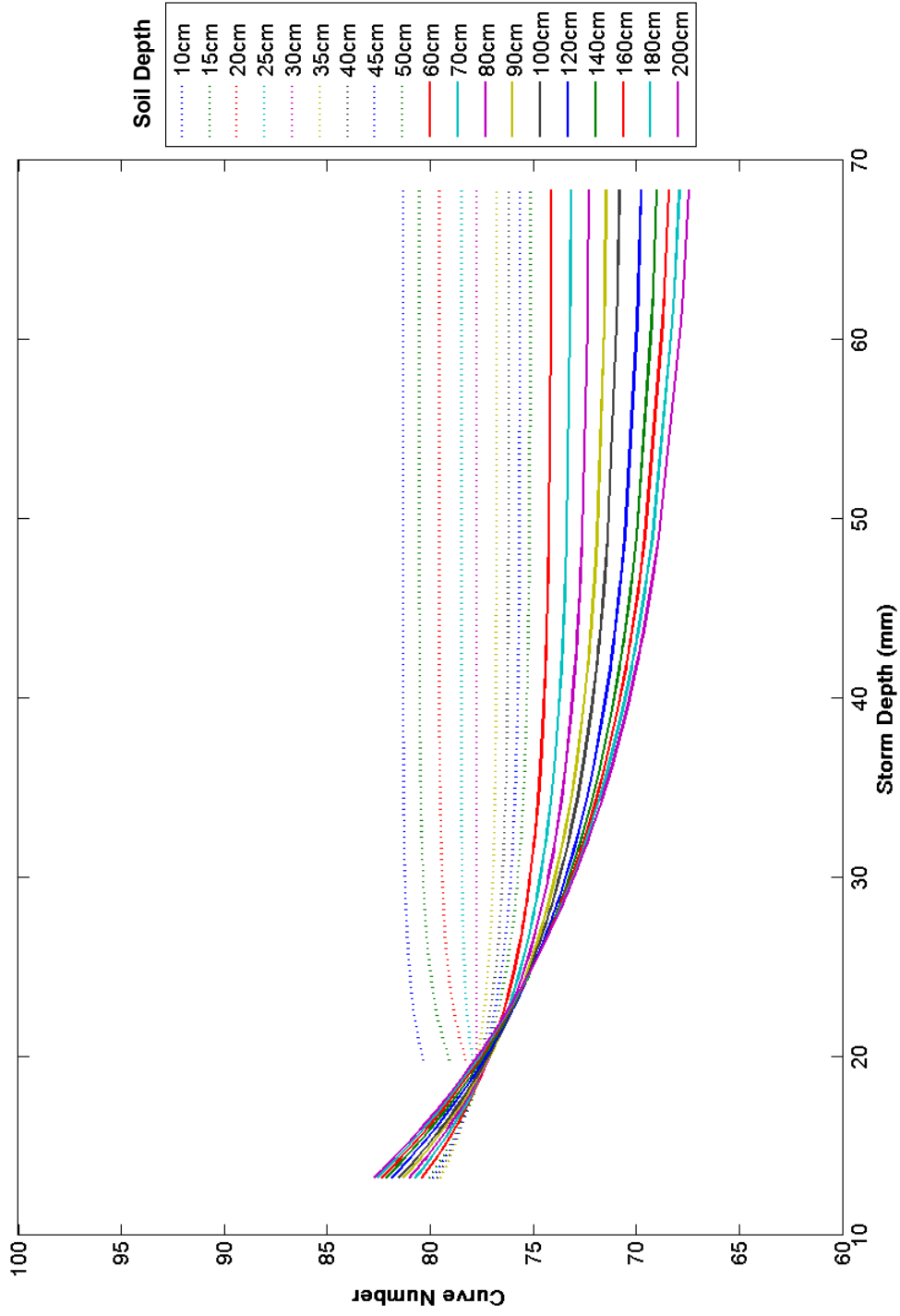


Figure 35. Asymptotic Method, CN Fit Curves vs. Storm Depth, Silt Loam, $\lambda = 0.2$, Beckley

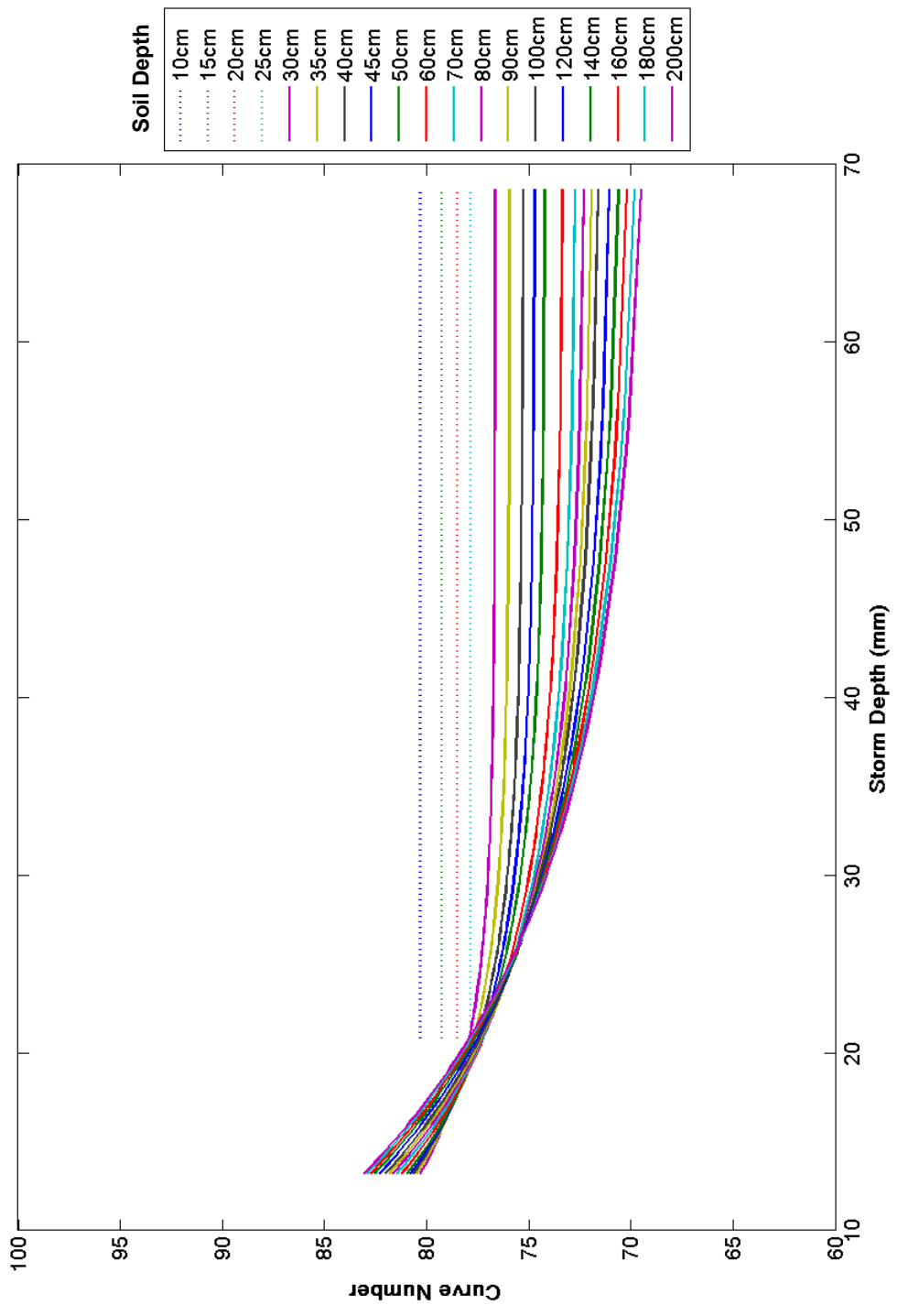


Figure 36. Asymptotic Method, CN Fit Curves vs. Storm Depth, Silt Loam $\lambda = 0.2$, Elkins

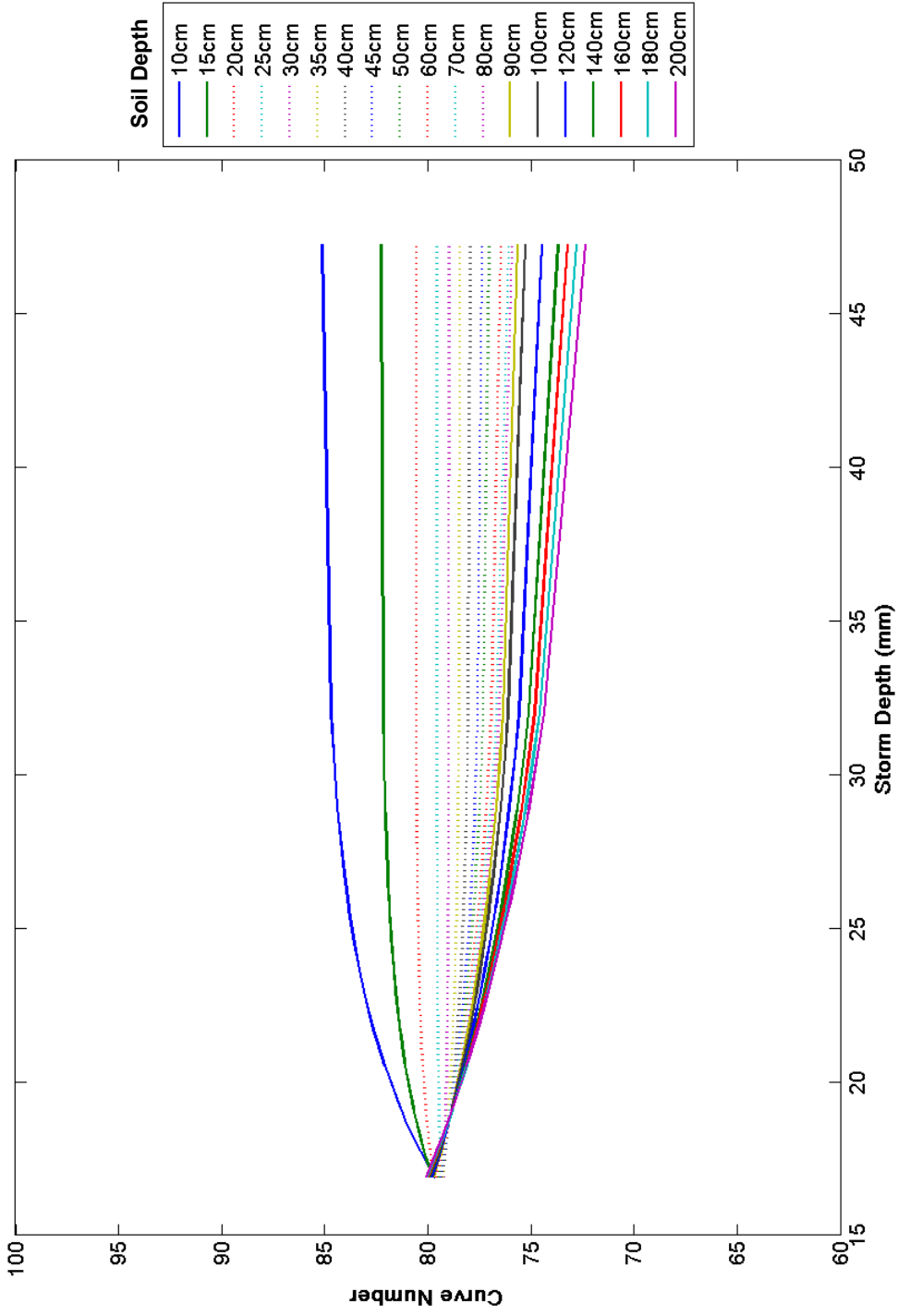


Figure 37. Asymptotic Method, CN Fit Curves vs. Storm Depth, Silt Loam, $\lambda = 0.2$, Dunlow

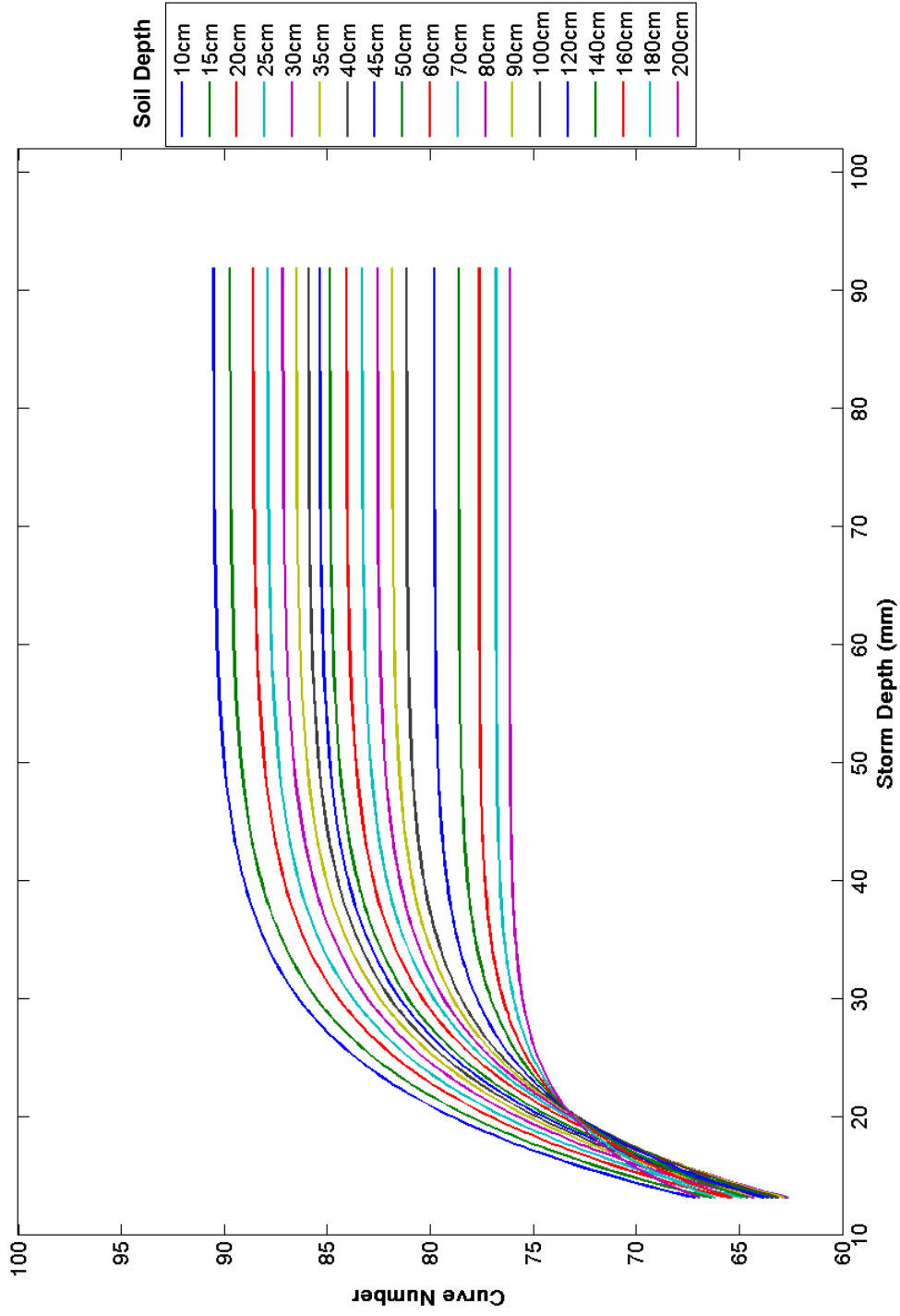


Figure 38. Asymptotic Method, CN Fit Curves vs. Storm Depth, Clay Loam, $\lambda = 0.05$, Terra Alta

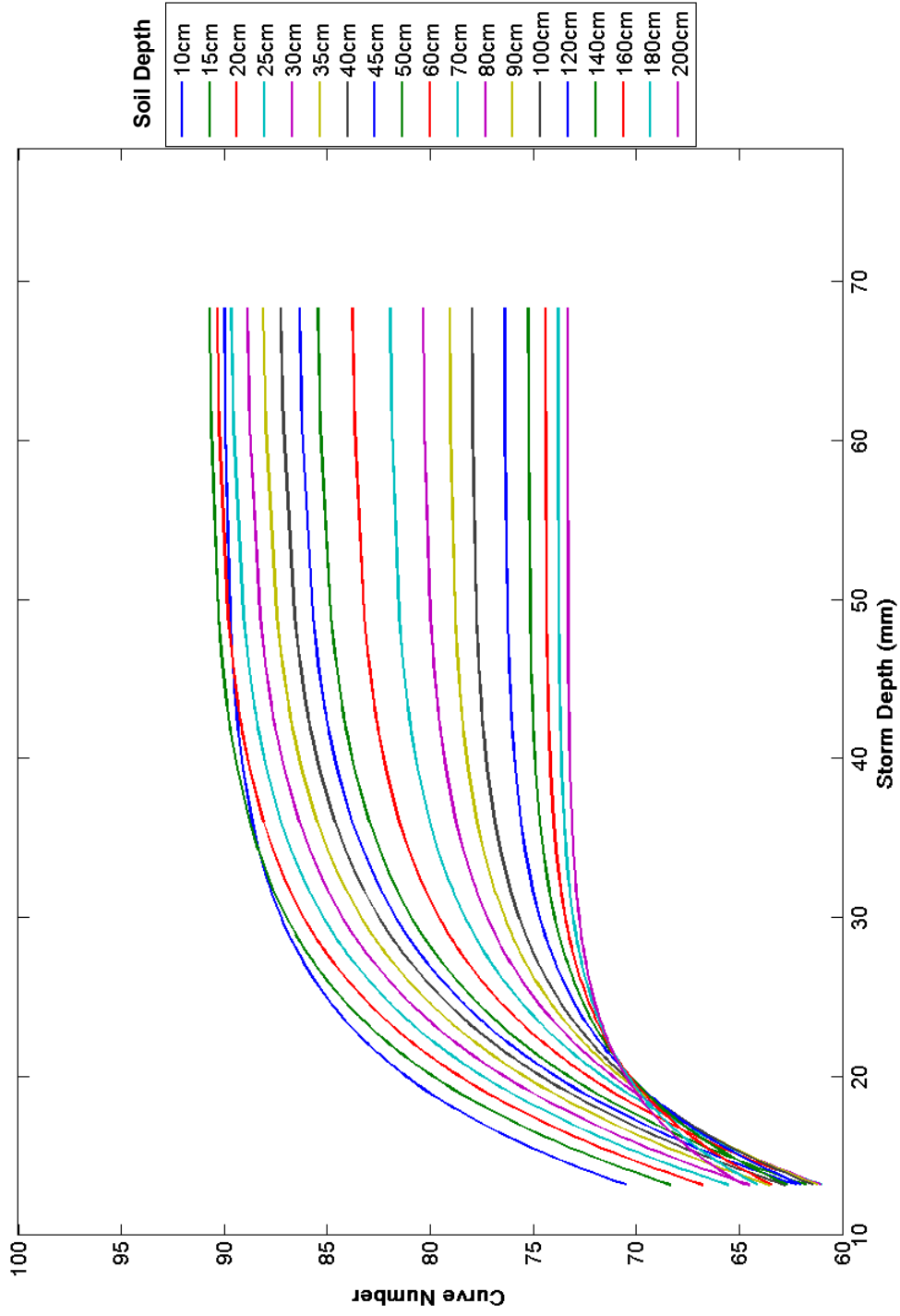


Figure 39. Asymptotic Method, CN Fit Curves vs. Storm Depth, Clay Loam, $\lambda = 0.05$, Beckley

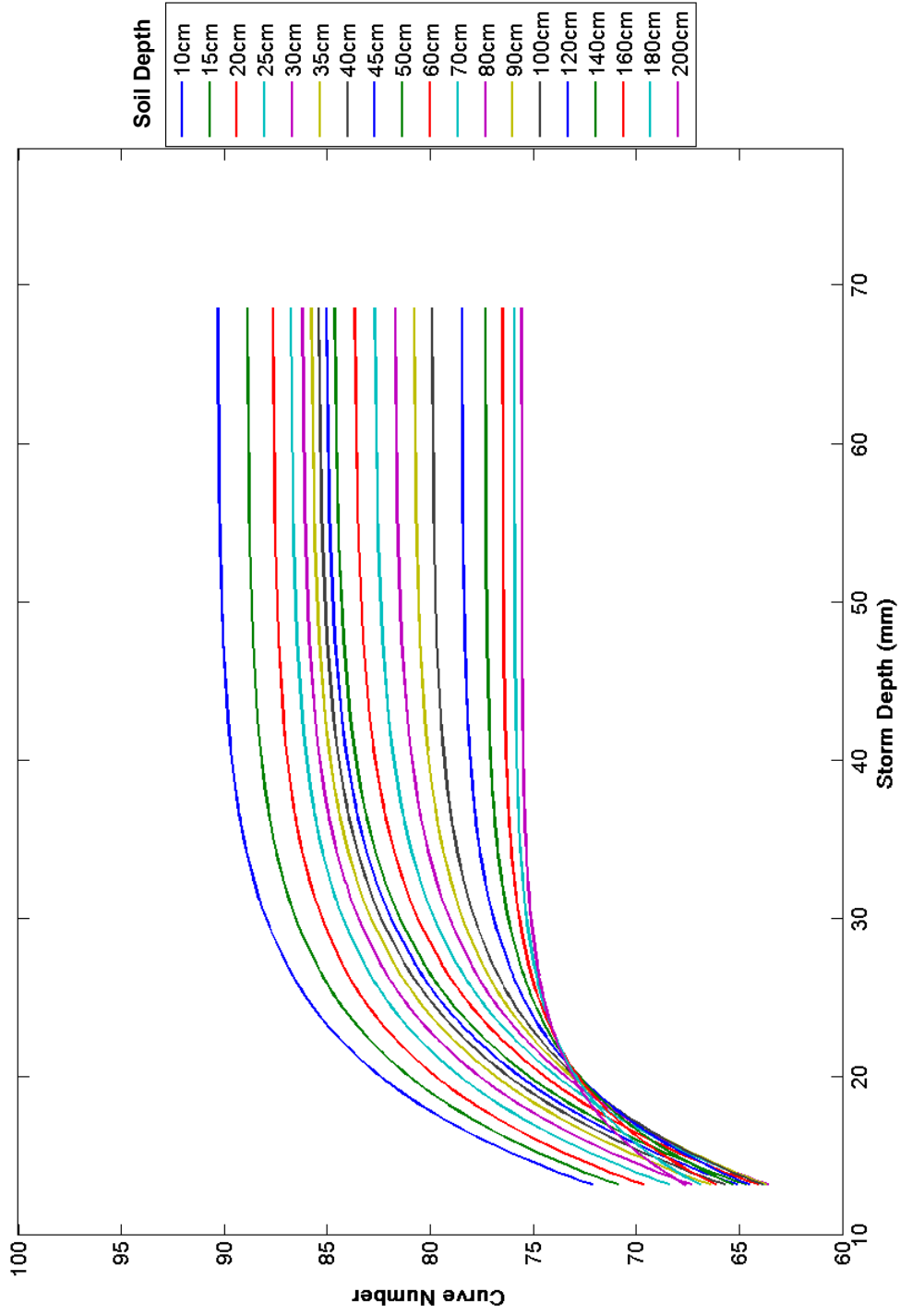


Figure 40. Asymptotic Method, CN Fit Curves vs. Storm Depth, Clay Loam, $\lambda = 0.05$, Elkins

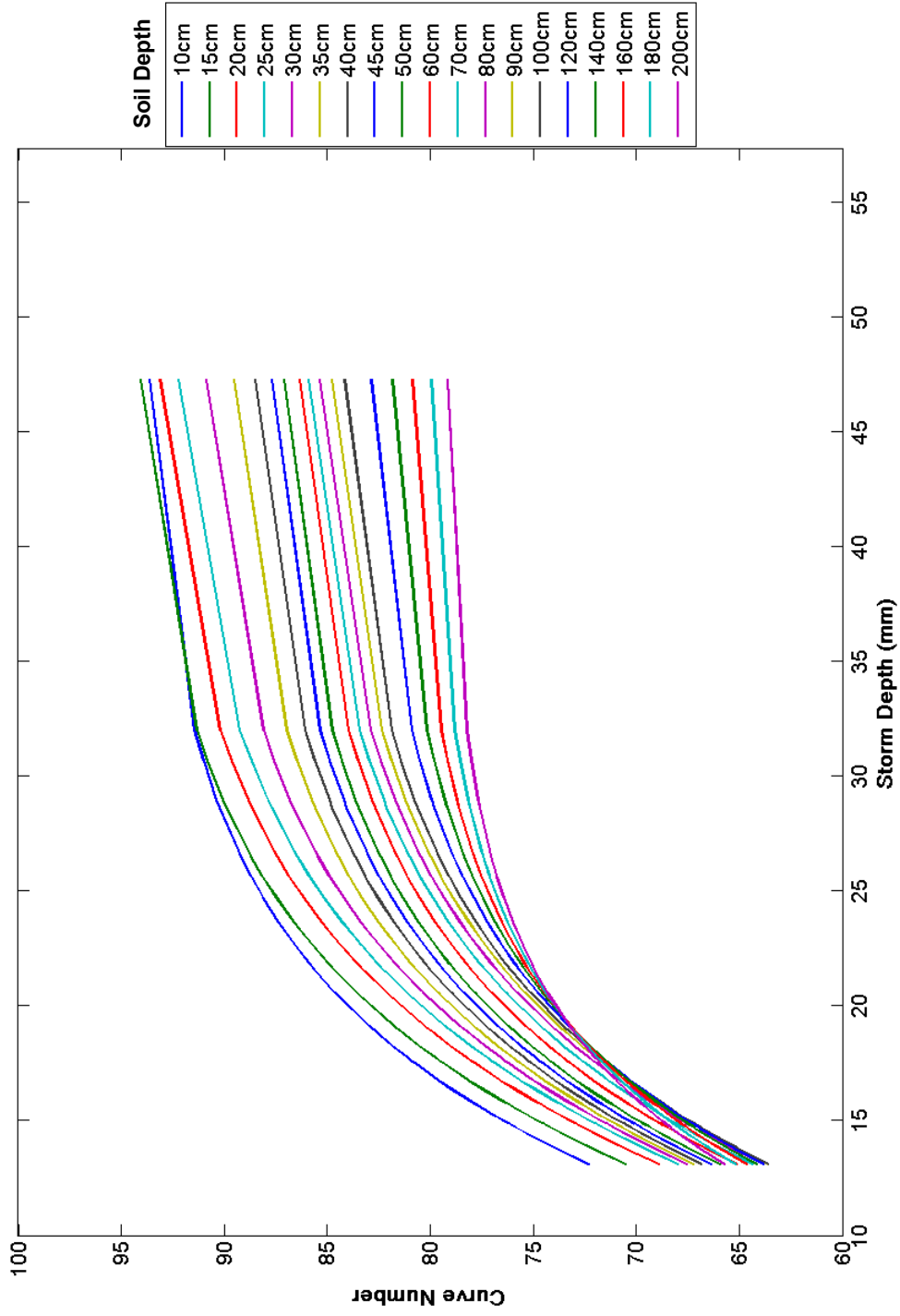


Figure 41. Asymptotic Method, CN Fit Curves vs. Storm Depth, Clay Loam, $\lambda = 0.05$, Dunlow

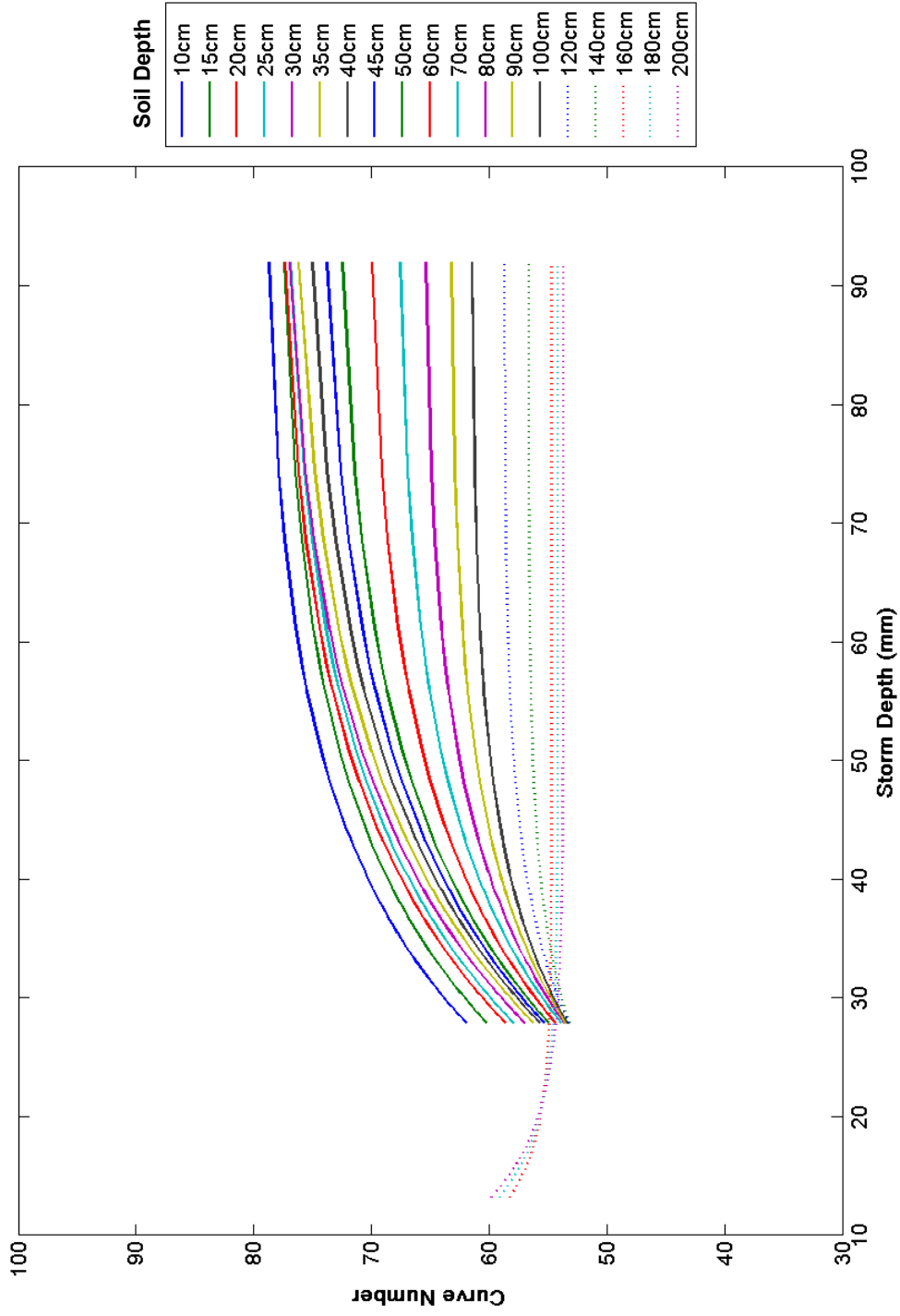


Figure 42. Asymptotic Method, CN Fit Curves vs. Storm Depth, Silt Loam, $\lambda = 0.05$, Terra Alta

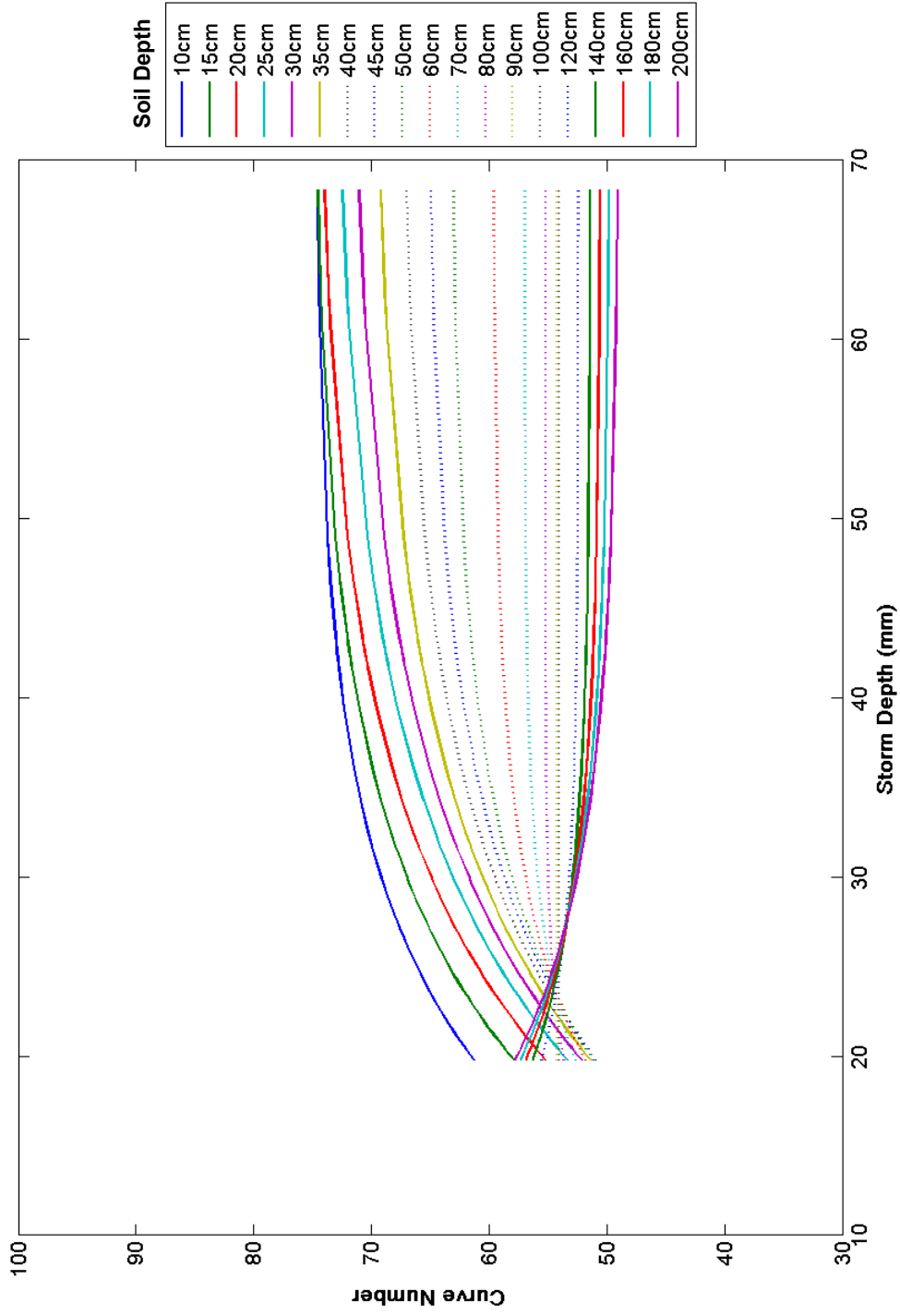


Figure 43. Asymptotic Method, CN Fit Curves vs. Storm Depth, Silt Loam, $\lambda = 0.05$, Beckley

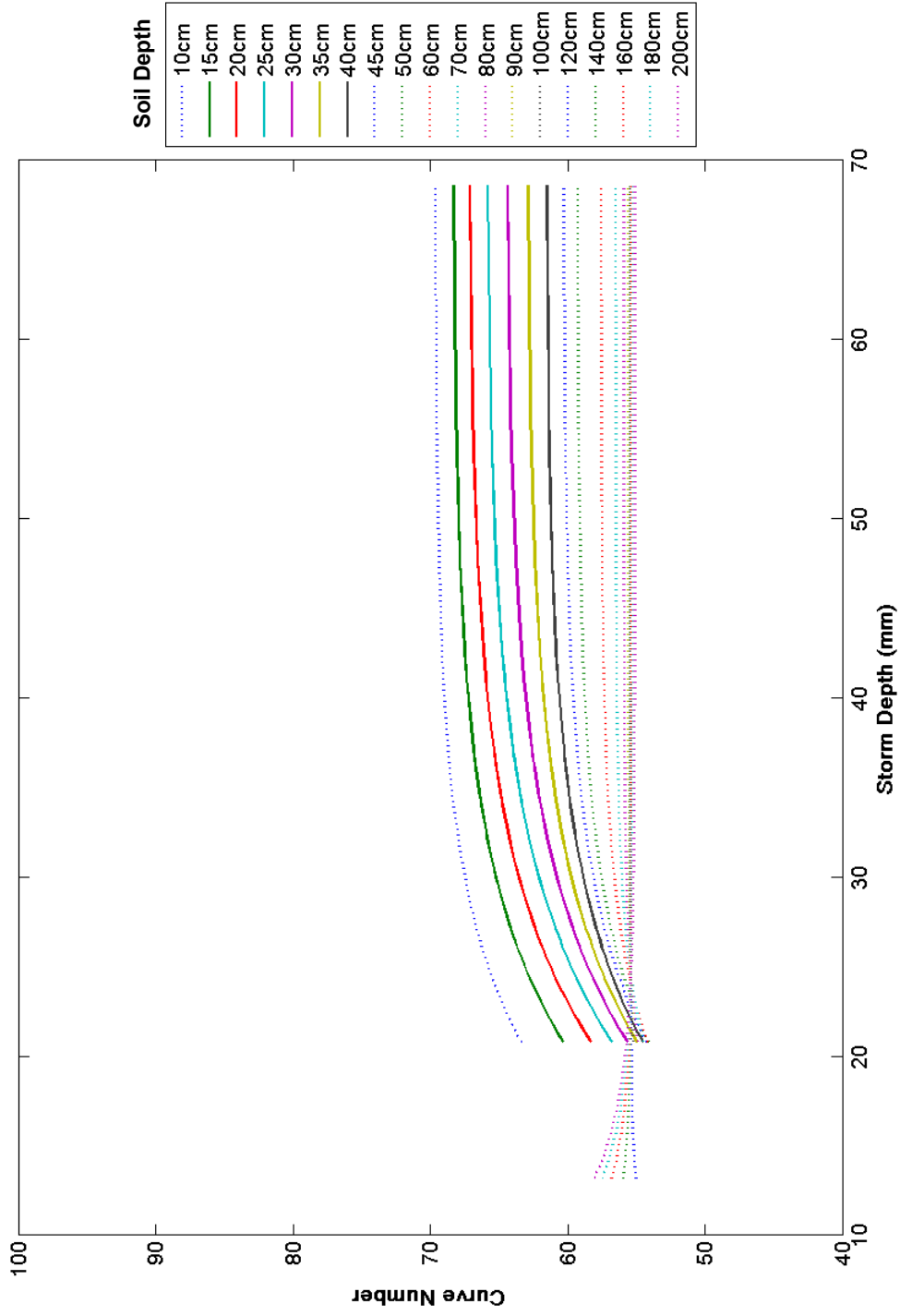


Figure 44. Asymptotic Method, CN Fit Curves vs. Storm Depth, Silt Loam, $\lambda = 0.05$, Elkins

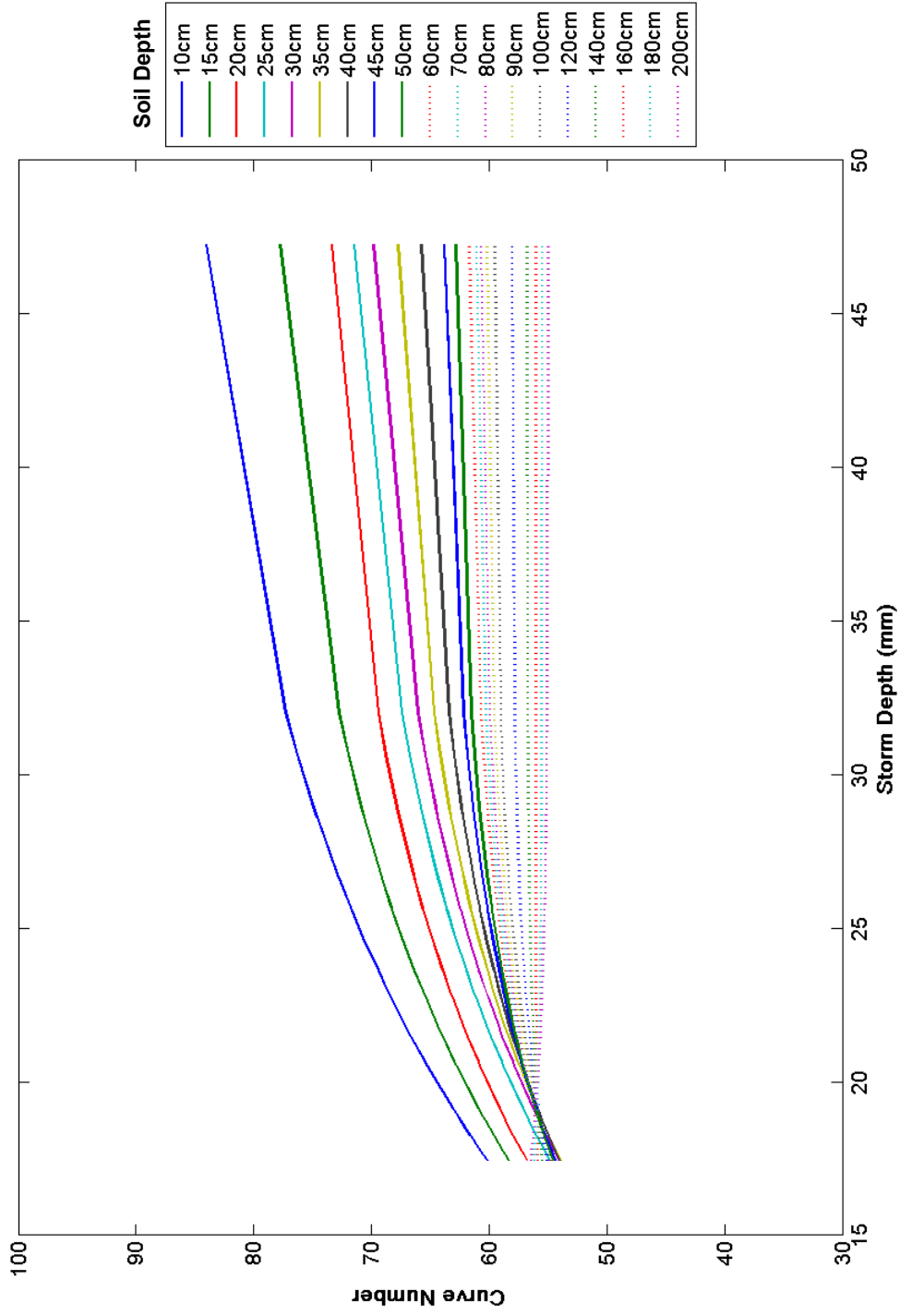


Figure 45. Asymptotic Method, CN Fit Curves vs. Storm Depth, Silt Loam, $\lambda = 0.05$, Dunlow

Table 4. Asymptotic Method Curve Fits, $\lambda = 0.2$ ('s' = standard, Eqn. 15; 'v' = violent, Eqn. 16)

Soil Depth (cm)	Terra Alta (163 storms)						Beckley (133 storms)					
	Clay Loam			Silt Loam			Clay Loam			Silt Loam		
	CN_{∞}	k	R^2	CN_{∞}	k	R^2	CN_{∞}	k	R^2	CN_{∞}	k	R^2
10	90.71	4.75	0.82v	82.04	2.65	0.71v	90.63	5.24	0.71v	81.32	5.66	0.03v
15	90.17	4.82	0.82v	81.13	2.66	0.57v	90.56	4.97	0.82v	80.52	5.14	0.05v
20	89.46	4.94	0.81v	80.76	2.58	0.55v	90.09	4.92	0.87v	79.57	5.28	0.03v
25	89.02	5.02	0.80v	80.44	2.57	0.57v	89.55	4.95	0.89v	78.49	6.35	0.00v
30	88.57	5.12	0.78v	80.26	2.52	0.61v	88.99	5.01	0.88v	77.73	44.86	0.00v
35	88.18	5.22	0.75v	79.69	2.57	0.63v	88.45	5.09	0.84v	76.77	4.13	0.22s
40	87.86	5.32	0.72v	79.01	2.68	0.58v	87.93	5.20	0.79v	76.18	3.69	0.31s
45	87.56	5.42	0.67v	78.33	2.83	0.48v	87.43	5.33	0.75v	75.63	3.37	0.40s
50	87.30	5.52	0.63v	76.06	3.29	0.63s	86.97	5.47	0.69v	75.11	3.12	0.49s
60	86.87	5.68	0.59v	75.40	2.94	0.74s	86.20	5.78	0.58v	74.10	2.72	0.66s
70	86.53	5.85	0.57v	74.73	2.67	0.81s	85.51	6.19	0.43v	73.12	2.43	0.81s
80	86.22	6.07	0.56v	74.08	2.45	0.87s	84.96	6.64	0.28v	72.22	2.21	0.89s
90	85.95	6.31	0.54v	73.38	2.26	0.90s	84.56	7.12	0.18v	71.33	2.03	0.92s
100	85.70	6.58	0.48v	72.75	2.11	0.91s	84.25	7.64	0.10v	70.64	1.91	0.93s
120	85.26	7.30	0.31v	71.58	1.88	0.93s	83.86	8.83	0.03v	69.48	1.73	0.92s
140	84.91	8.29	0.13v	70.51	1.71	0.94s	83.65	10.59	0.01v	68.60	1.62	0.91s
160	84.66	9.65	0.04v	69.57	1.59	0.95s	82.96	4.65	0.04s	67.90	1.53	0.90s
180	84.09	4.81	0.05s	68.68	1.49	0.96s	82.35	3.60	0.13s	67.27	1.47	0.90s
200	83.57	3.68	0.18s	67.90	1.41	0.96s	81.77	3.05	0.24s	66.68	1.41	0.89s
Soil Depth (cm)	Elkins (137 storms)						Dunlow (45 storms)					
	Clay Loam			Silt Loam			Clay Loam			Silt Loam		
	CN_{∞}	k	R^2	CN_{∞}	k	R^2	CN_{∞}	k	R^2	CN_{∞}	k	R^2
10	91.14	5.29	0.58v	80.30	12.65	0.00v	92.47	4.97	0.67v	85.15	4.11	0.58v
15	90.23	5.46	0.61v	79.25	12.70	0.00v	92.18	4.86	0.79v	82.27	5.24	0.50v
20	89.47	5.59	0.60v	78.49	14.66	0.00v	91.39	4.91	0.80v	80.56	6.84	0.06v
25	88.87	5.65	0.61v	77.86	20.27	0.00v	90.77	4.99	0.83v	79.55	8.89	0.00v
30	88.45	5.67	0.62v	76.64	3.57	0.55s	90.10	5.17	0.84v	78.97	6.55	0.01s
35	88.12	5.67	0.61v	75.92	3.21	0.68s	89.47	5.37	0.82v	78.44	4.89	0.06s
40	87.86	5.67	0.61v	75.27	2.94	0.75s	88.98	5.52	0.80v	77.92	4.16	0.22s
45	87.60	5.69	0.61v	74.69	2.74	0.78s	88.57	5.64	0.79v	77.33	3.58	0.39s
50	87.36	5.73	0.59v	74.18	2.58	0.83s	88.25	5.73	0.78v	76.96	3.33	0.38s
60	86.87	5.89	0.56v	73.29	2.34	0.88s	87.81	5.81	0.76v	76.37	2.99	0.30s
70	86.42	6.11	0.53v	72.66	2.19	0.91s	87.50	5.83	0.82v	75.94	2.78	0.26s
80	86.01	6.39	0.48v	72.20	2.08	0.93s	87.21	5.89	0.83v	75.76	2.74	0.35s
90	85.67	6.71	0.39v	71.80	1.99	0.93s	86.96	5.98	0.84v	75.45	2.64	0.50s
100	85.38	7.09	0.30v	71.44	1.91	0.92s	86.75	6.10	0.83v	75.04	2.51	0.69s
120	84.96	8.01	0.14v	70.82	1.80	0.91s	86.36	6.42	0.71v	74.09	2.27	0.89s
140	84.71	9.23	0.04v	70.30	1.71	0.92s	86.08	6.75	0.53v	73.10	2.06	0.95s
160	84.21	5.42	0.01s	69.82	1.63	0.91s	85.84	7.15	0.33v	72.48	1.96	0.95s
180	83.62	3.83	0.10s	69.36	1.57	0.91s	85.64	7.65	0.18v	71.90	1.87	0.97s
200	83.19	3.26	0.19s	68.99	1.52	0.90s	85.46	8.28	0.10v	71.32	1.78	0.98s

Table 5. Asymptotic Method Curve Fits, $\lambda = 0.05$ ('s' = standard, Eqn. 15; 'v' = violent, Eqn. 16)

Soil Depth (cm)	Terra Alta (163 storms)						Beckley (133 storms)					
	Clay Loam			Silt Loam			Clay Loam			Silt Loam		
	CN_{∞}	k	R^2	CN_{∞}	k	R^2	CN_{∞}	k	R^2	CN_{∞}	k	R^2
10	90.53	2.60	0.75v	79.26	1.38	0.85v	90.02	2.94	0.68v	74.79	2.19	0.56v
15	89.73	2.58	0.77v	78.08	1.34	0.79v	90.77	2.69	0.78v	74.98	1.89	0.67v
20	88.60	2.58	0.78v	78.18	1.26	0.79v	90.42	2.58	0.83v	74.69	1.72	0.68v
25	87.90	2.58	0.80v	77.83	1.24	0.78v	89.75	2.52	0.86v	73.24	1.67	0.63v
30	87.17	2.57	0.82v	78.05	1.19	0.80v	88.99	2.48	0.88v	71.84	1.66	0.59v
35	86.51	2.58	0.82v	77.27	1.18	0.83v	88.22	2.45	0.90v	69.88	1.70	0.55v
40	85.92	2.59	0.82v	75.98	1.20	0.85v	87.37	2.44	0.90v	67.51	1.81	0.48v
45	85.37	2.61	0.82v	74.57	1.23	0.85v	86.46	2.45	0.91v	65.30	1.94	0.42v
50	84.88	2.62	0.81v	73.16	1.27	0.84v	85.55	2.46	0.90v	63.23	2.12	0.35v
60	84.07	2.64	0.82v	70.37	1.35	0.77v	83.85	2.52	0.90v	59.62	2.61	0.22v
70	83.32	2.68	0.83v	67.77	1.45	0.70v	82.01	2.62	0.88v	56.97	3.32	0.12v
80	82.56	2.74	0.84v	65.50	1.56	0.68v	80.38	2.74	0.85v	55.22	4.54	0.03v
90	81.84	2.81	0.85v	63.25	1.70	0.61v	79.09	2.86	0.82v	54.24	46.87	0.00v
100	81.15	2.89	0.85v	61.45	1.85	0.52v	77.99	2.98	0.79v	54.04	48.30	0.00v
120	79.79	3.09	0.84v	58.69	2.16	0.39v	76.40	3.21	0.75v	52.43	3.50	0.38s
140	78.60	3.32	0.81v	56.58	2.57	0.24v	75.24	3.45	0.71v	51.40	2.95	0.50s
160	77.62	3.55	0.77v	54.68	4.89	0.23s	74.40	3.68	0.67v	50.54	2.64	0.58s
180	76.80	3.80	0.72v	54.14	4.28	0.39s	73.78	3.91	0.64v	49.74	2.42	0.59s
200	76.13	4.06	0.66v	53.69	3.90	0.48s	73.32	4.13	0.61v	48.96	2.26	0.61s
Soil Depth (cm)	Elkins (137 storms)						Dunlow (45 storms)					
	Clay Loam			Silt Loam			Clay Loam			Silt Loam		
	CN_{∞}	k	R^2	CN_{∞}	k	R^2	CN_{∞}	k	R^2	CN_{∞}	k	R^2
10	90.33	3.08	0.61v	69.63	2.94	0.44v	94.09	2.83	0.65v	87.60	1.69	0.87v
15	88.88	3.07	0.65v	68.33	2.61	0.54v	94.74	2.64	0.75v	79.88	1.91	0.95v
20	87.65	3.04	0.68v	67.18	2.46	0.64v	93.92	2.56	0.80v	74.92	2.07	0.85v
25	86.78	2.98	0.72v	65.90	2.41	0.63v	93.04	2.54	0.84v	73.00	2.03	0.74v
30	86.22	2.92	0.75v	64.44	2.42	0.60v	91.62	2.58	0.86v	70.96	2.10	0.72v
35	85.79	2.86	0.77v	62.93	2.52	0.57v	90.18	2.65	0.85v	68.32	2.29	0.63v
40	85.45	2.81	0.79v	61.54	2.65	0.51v	89.08	2.68	0.86v	66.02	2.51	0.62v
45	85.07	2.79	0.79v	60.30	2.81	0.41v	88.27	2.70	0.87v	63.93	2.79	0.57v
50	84.66	2.77	0.80v	59.28	2.97	0.37v	87.65	2.70	0.88v	63.00	2.94	0.51v
60	83.69	2.79	0.81v	57.56	3.45	0.25v	86.94	2.68	0.89v	61.85	3.15	0.36v
70	82.72	2.83	0.82v	56.48	4.08	0.14v	86.55	2.63	0.91v	61.27	3.27	0.27v
80	81.73	2.90	0.82v	55.88	4.92	0.05v	86.01	2.63	0.92v	60.89	3.33	0.32v
90	80.79	2.98	0.81v	55.56	5.90	0.02v	85.40	2.64	0.92v	60.31	3.47	0.38v
100	79.92	3.09	0.80v	55.43	6.83	0.01v	84.73	2.69	0.92v	59.57	3.68	0.46v
120	78.46	3.32	0.77v	55.37	9.65	0.01s	83.30	2.81	0.91v	58.04	4.34	0.44v
140	77.33	3.58	0.71v	55.41	8.50	0.00s	82.16	2.94	0.89v	56.71	5.64	0.15v
160	76.48	3.84	0.64v	55.32	6.59	0.04s	81.11	3.08	0.85v	55.97	9.78	0.00v
180	75.91	4.10	0.58v	55.20	5.75	0.09s	80.13	3.25	0.80v	55.45	5.82	0.09s
200	75.55	4.33	0.52v	55.09	5.25	0.13s	79.27	3.42	0.76v	54.92	4.83	0.28s

The results of the Asymptotic Method, included as Figures 30-45 and Tables 4-5, also indicate a dependence of the CN on soil type, soil depth, storm depth, and storm distribution. As in the Cyclic Method, the CN values are higher for the Clay Loam soil than the Silt Loam soil. The CN's also generally decrease with increasing soil depth. The majority of the curve fits are classified as violent, with the exception of the Silt Loam soil where $\lambda = 0.2$. The least accurate curve fits (R^2 values in red, see Tables 4-5) typically arise in the transition zones between violent and standard behavior.

The violent behavior tends to be most prevalent at lesser soil depths, transitioning toward standard or complacent behavior at greater soil depths. This trend was also seen in the Cyclic Method (Figures 12-27). Complacent behavior was often found to precede violent behavior at low storm depths. Following the work of Hawkins (1993), the complacent data points that plotted below a selected threshold rainfall depth were ignored in these cases. Figure 34 is an example (Terra Alta gage, Silt Loam, $\lambda = 0.2$) where the complacent data points below a storm depth of 28 mm were ignored for the 10-45 cm soil depths. Figures 46-64 are the individual curve fits comprising Figure 34. It should be noted that in some cases a standard fit (Equation 15) resulted in a high R^2 value although the data may possibly be more accurately characterized as complacent. This occurred most often in the greater soil depths. An example is shown in Figure 64.

The scatter of the data points in Figures 46-64 demonstrates the considerable variation of the CN with storm depth, which can be compared to that of the Type II and WDM Triangular storm distributions of the Cyclic Method (Figures 18 and 19). This scatter appeared in the plots of the individual curve fits from all gages. It may be explained by noting that according to Figure 11, the storms recorded at the Terra Alta, Beckley, Elkins, and Dunlow gages consist of relatively high intensity hourly rainfall intervals comparable to the Type II and WDM Triangular storm distributions.

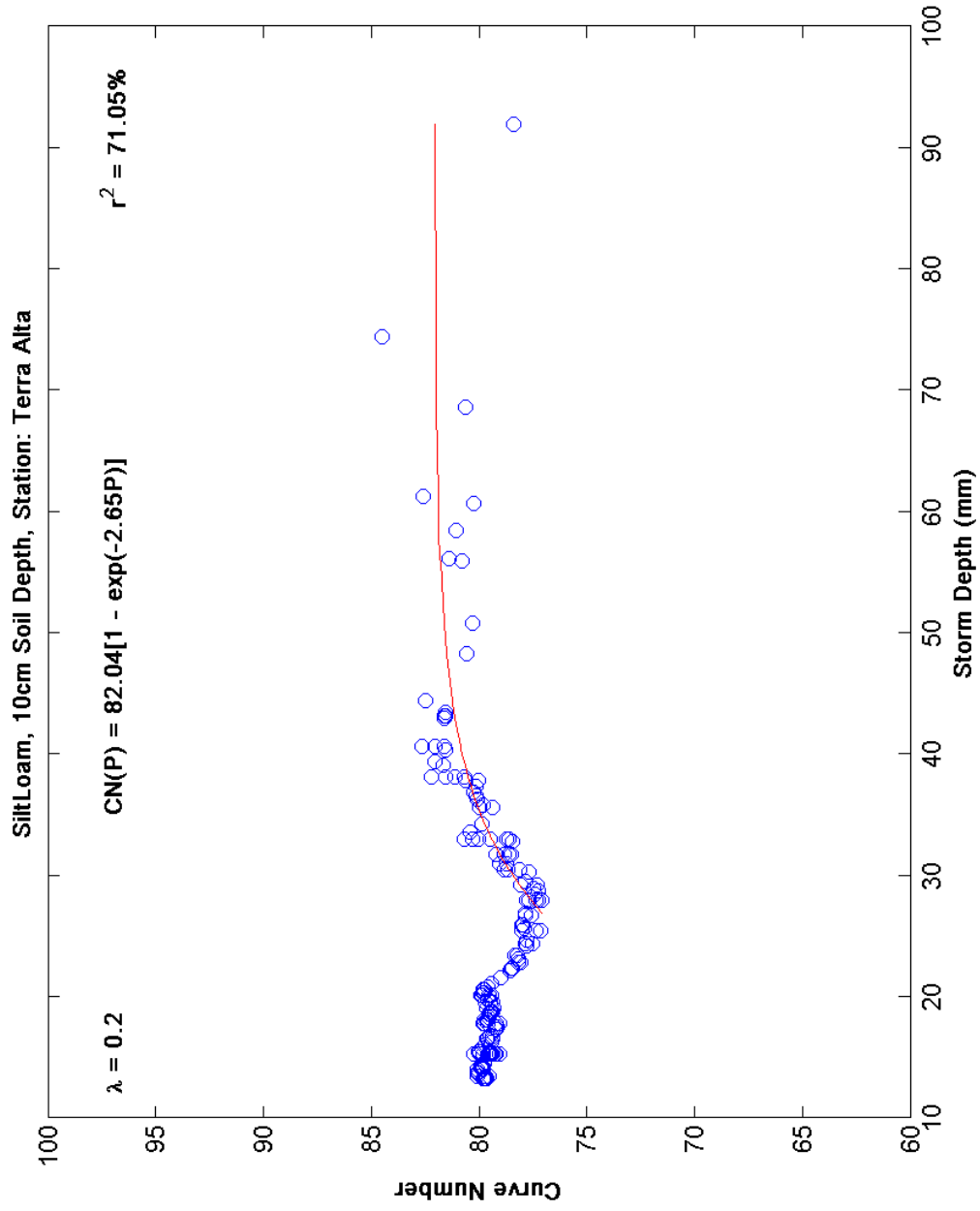


Figure 46. Asymptotic Method, CN vs. Storm Depth, Terra Alta, 10 cm Soil Depth

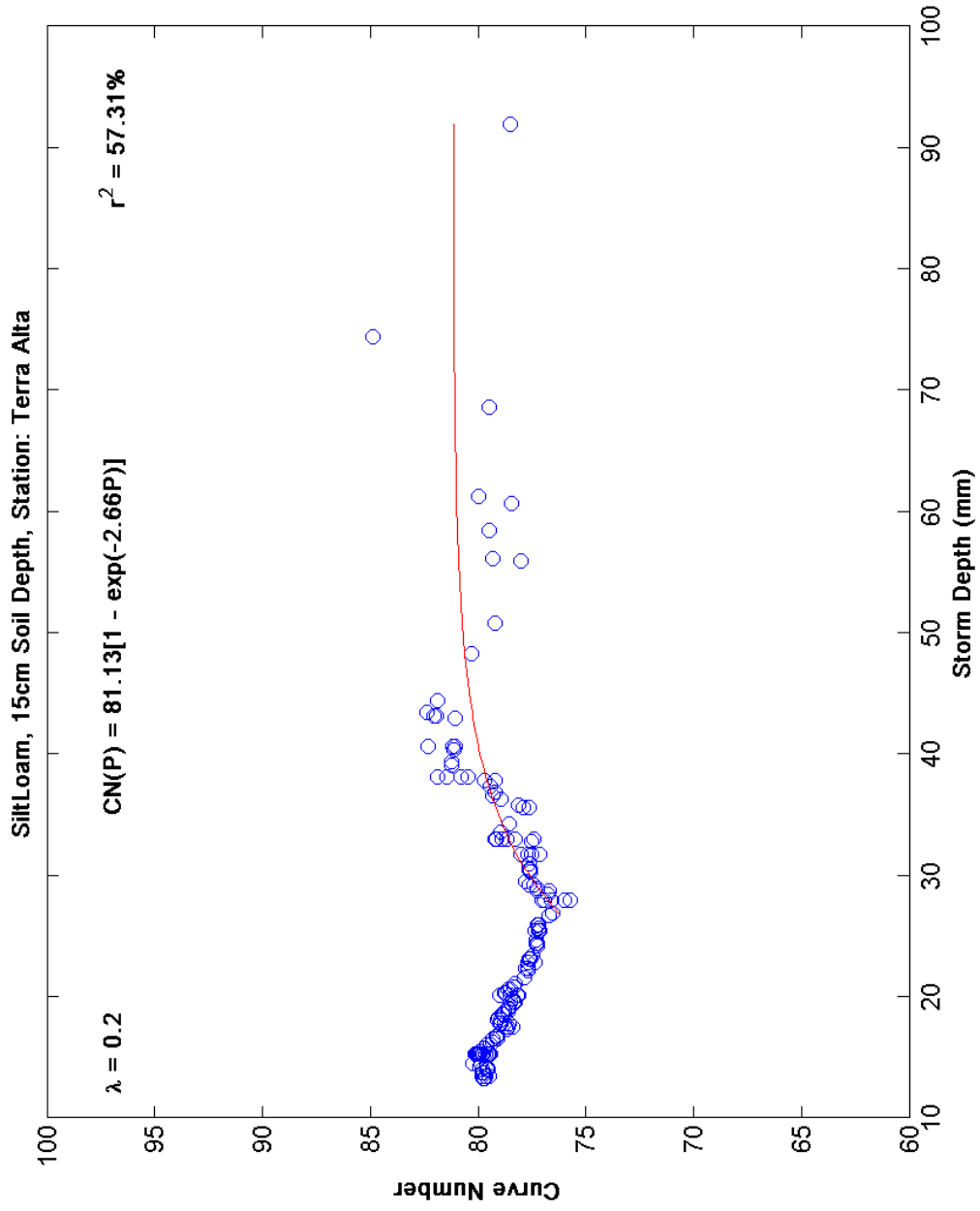


Figure 47. Asymptotic Method, CN vs. Storm Depth, Terra Alta, 15 cm Soil Depth

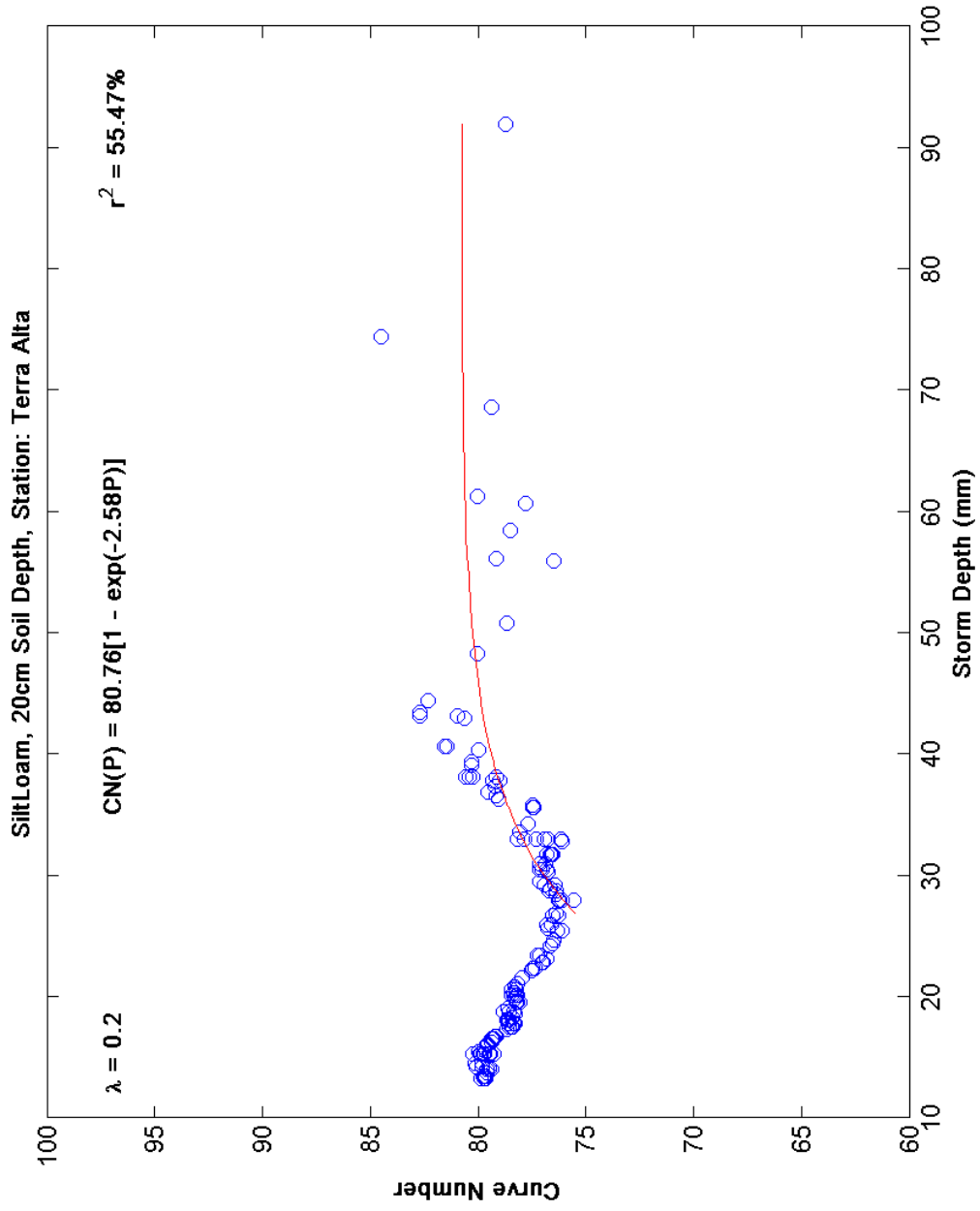


Figure 48. Asymptotic Method, CN vs. Storm Depth, Terra Alta, 20 cm Soil Depth

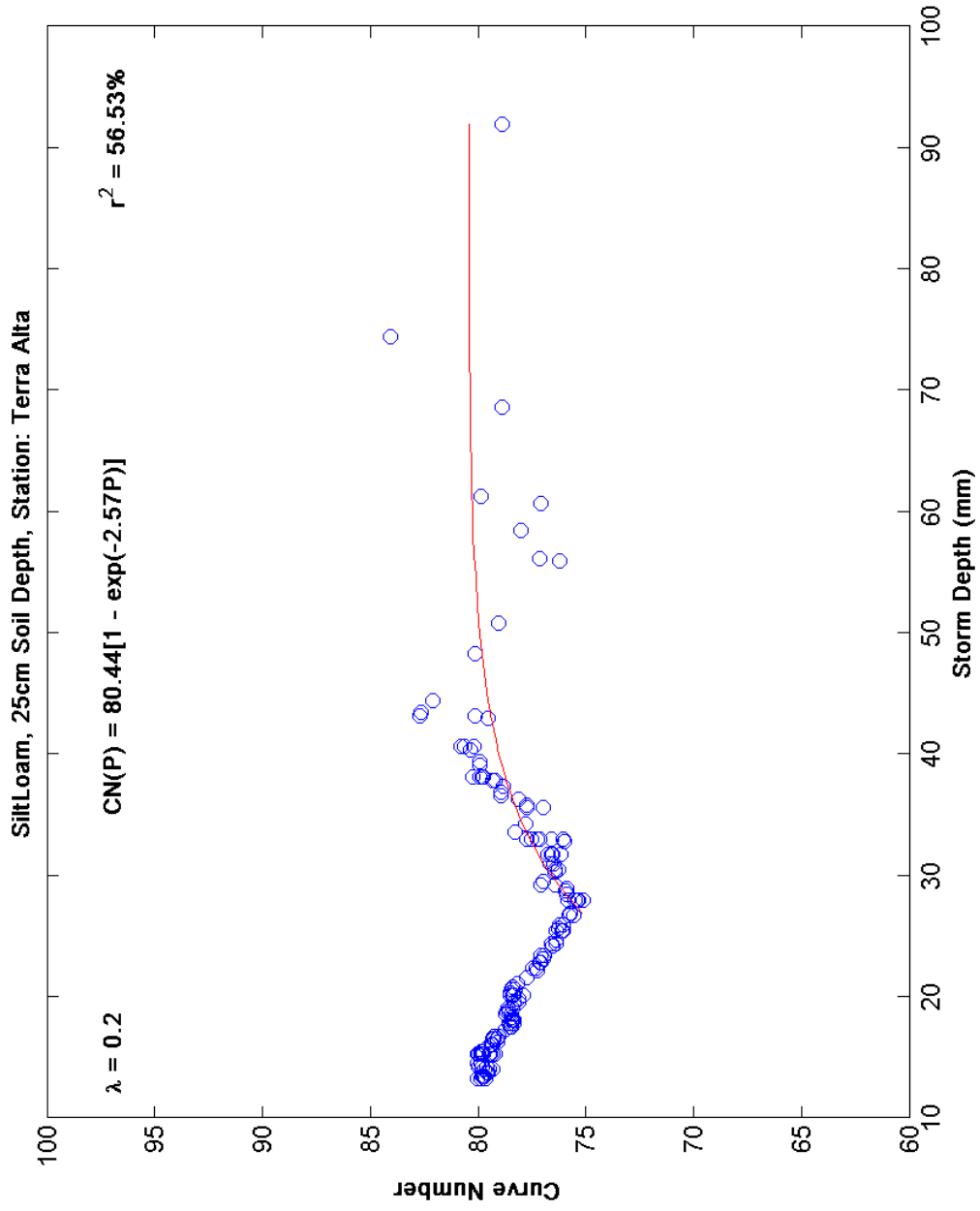


Figure 49. Asymptotic Method, CN vs. Storm Depth, Terra Alta, 25 cm Soil Depth

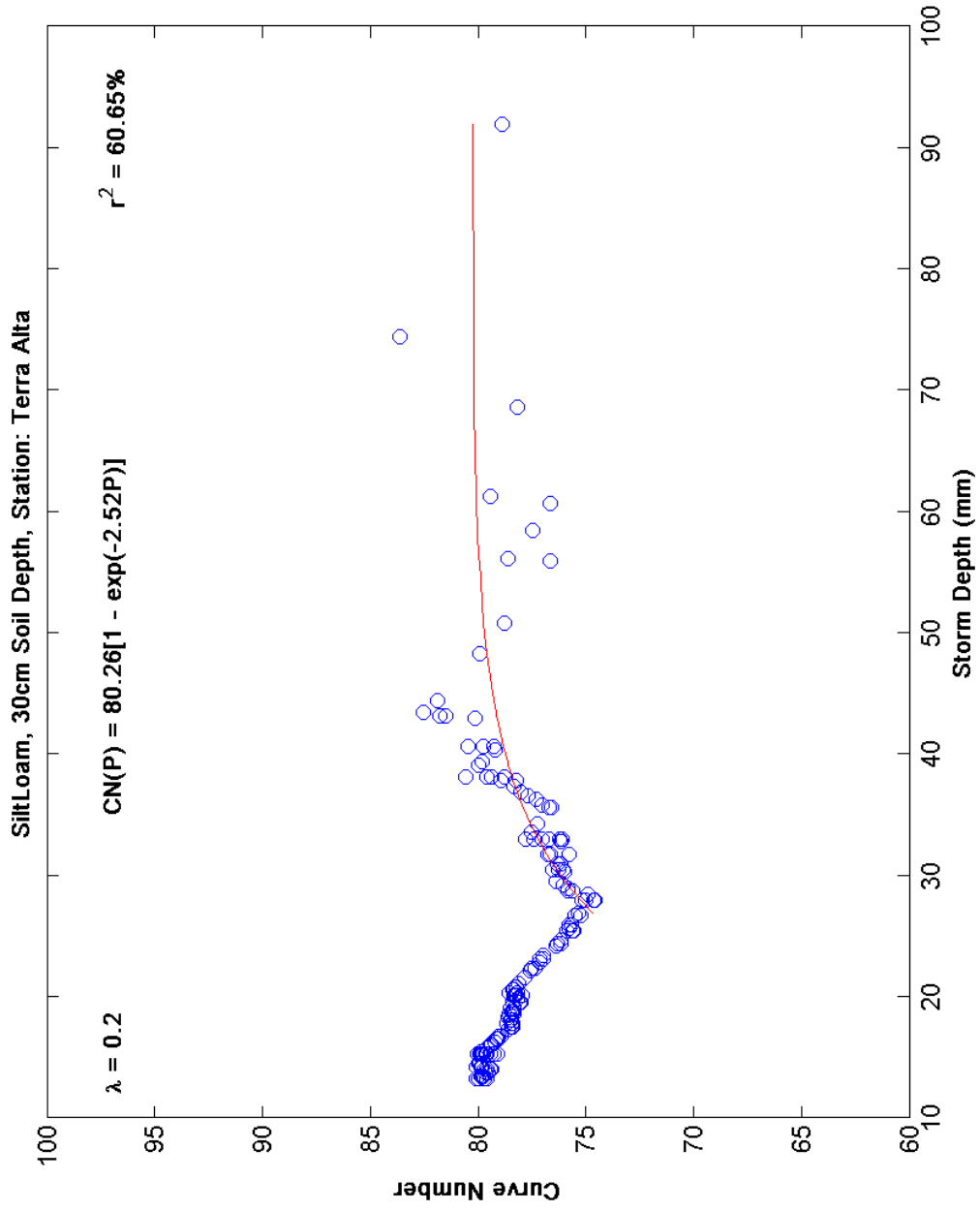


Figure 50. Asymptotic Method, CN vs. Storm Depth, Terra Alta, 30 cm Soil Depth

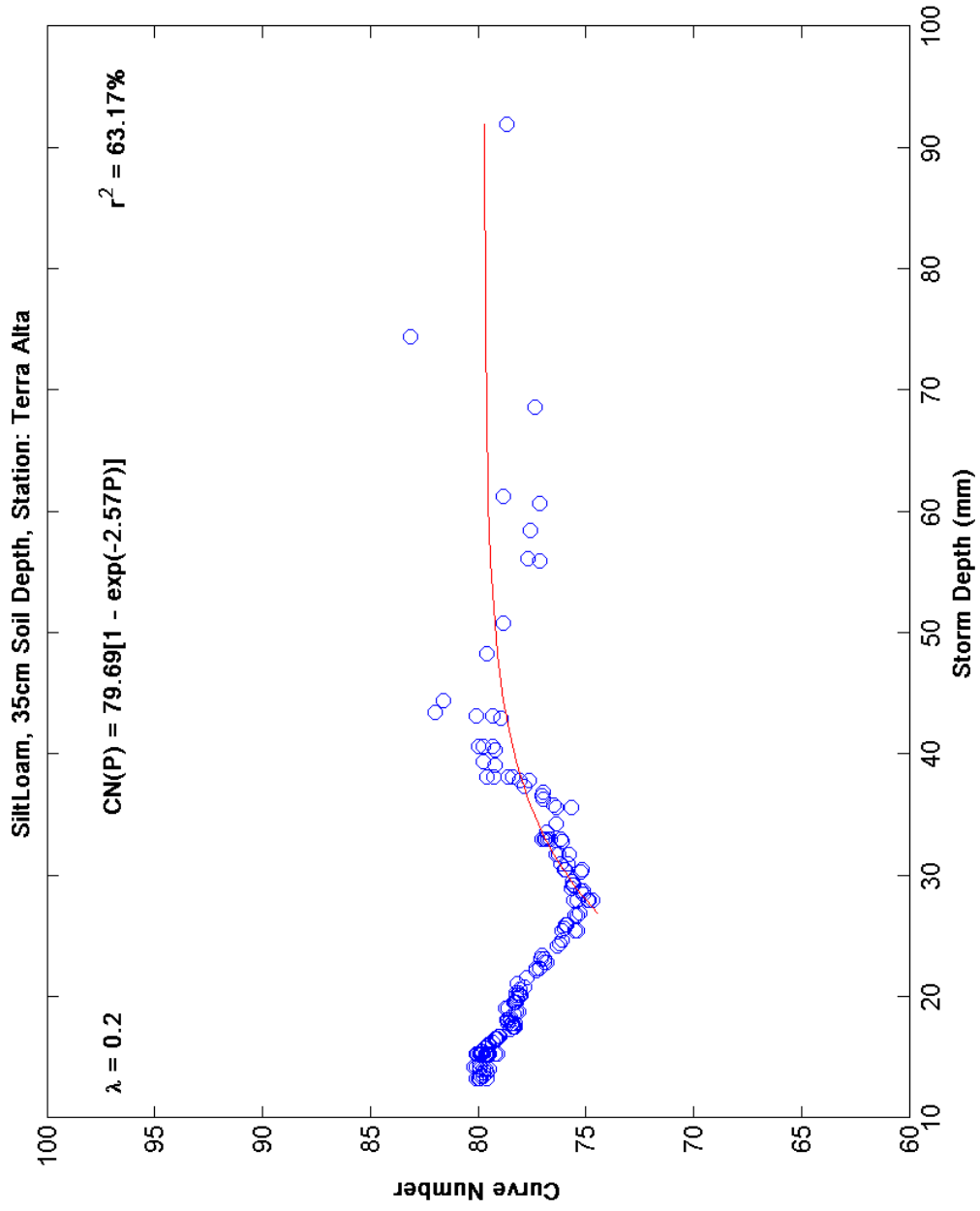


Figure 51. Asymptotic Method, CN vs. Storm Depth, Terra Alta, 35 cm Soil Depth

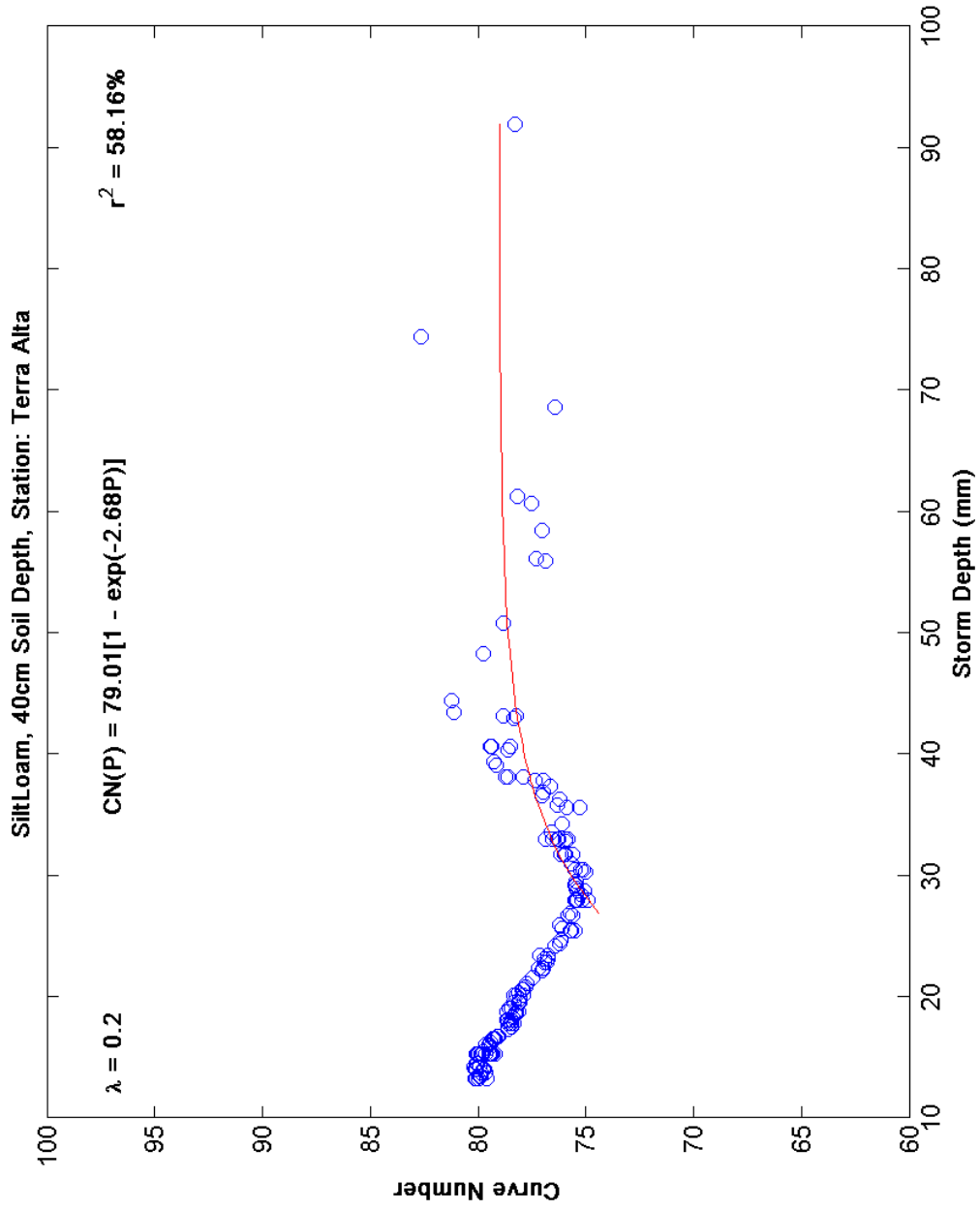


Figure 52. Asymptotic Method, CN vs. Storm Depth, Terra Alta, 40 cm Soil Depth

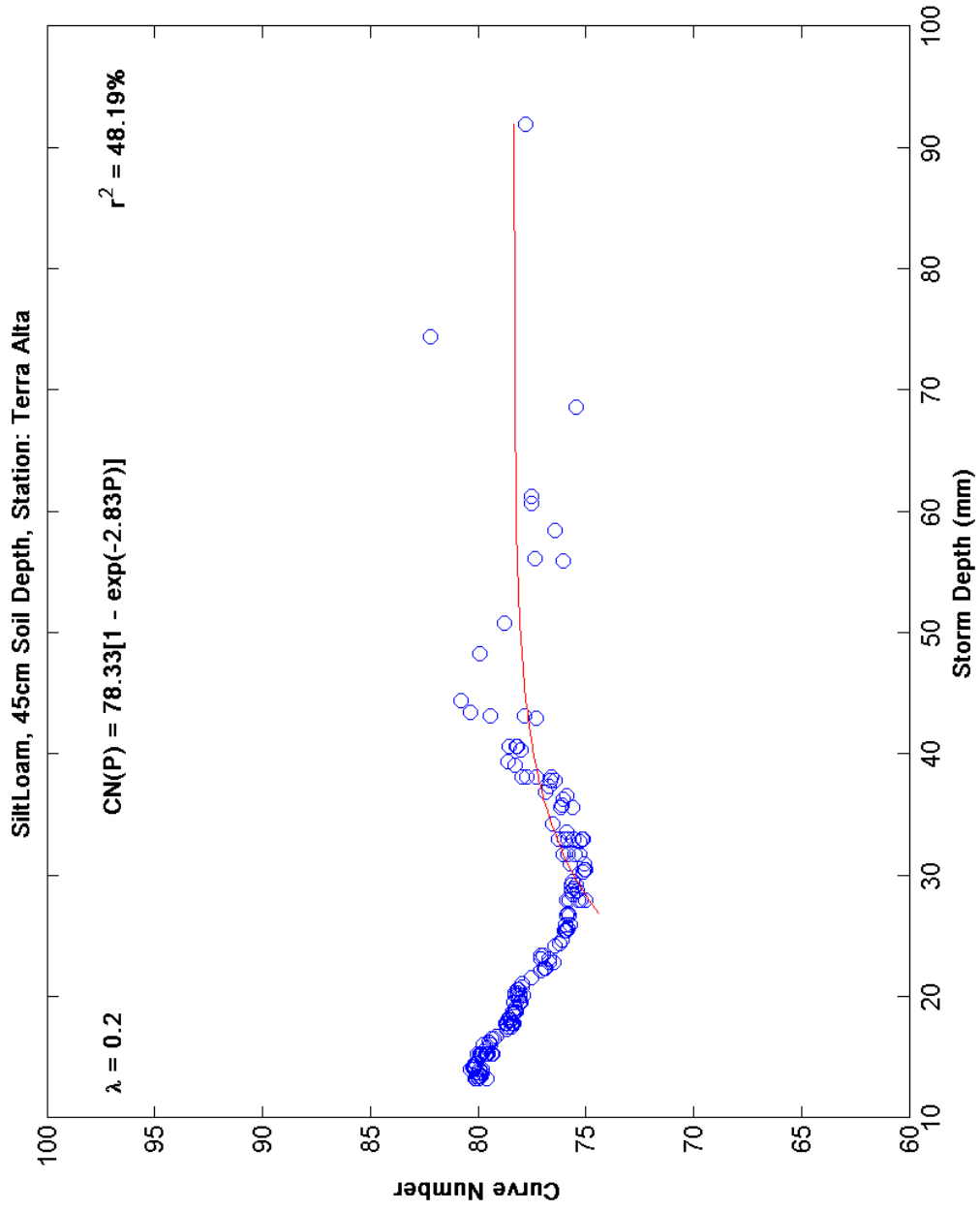


Figure 53. Asymptotic Method, CN vs. Storm Depth, Terra Alta, 45 cm Soil Depth

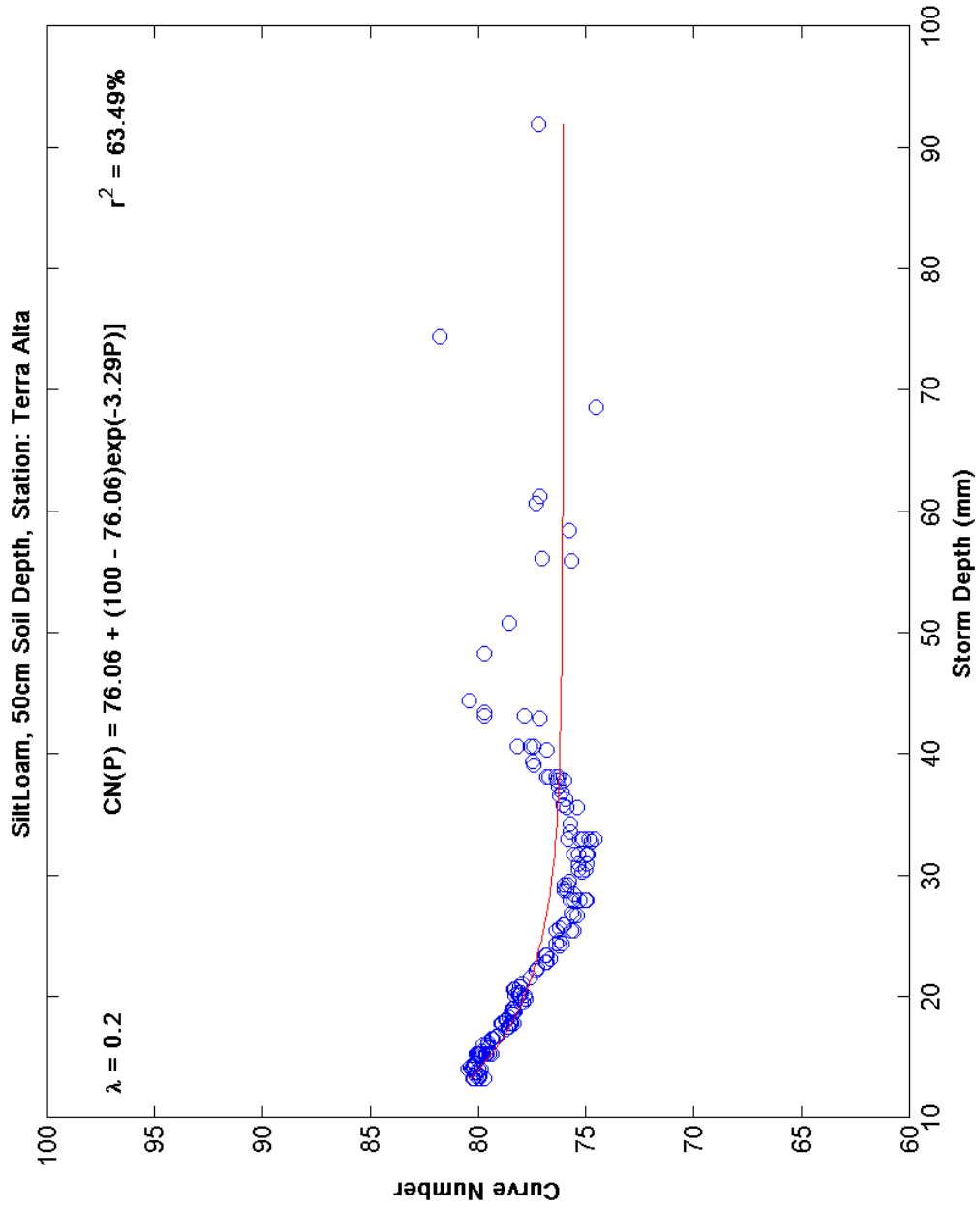


Figure 54. Asymptotic Method, CN vs. Storm Depth, Terra Alta, 50 cm Soil Depth

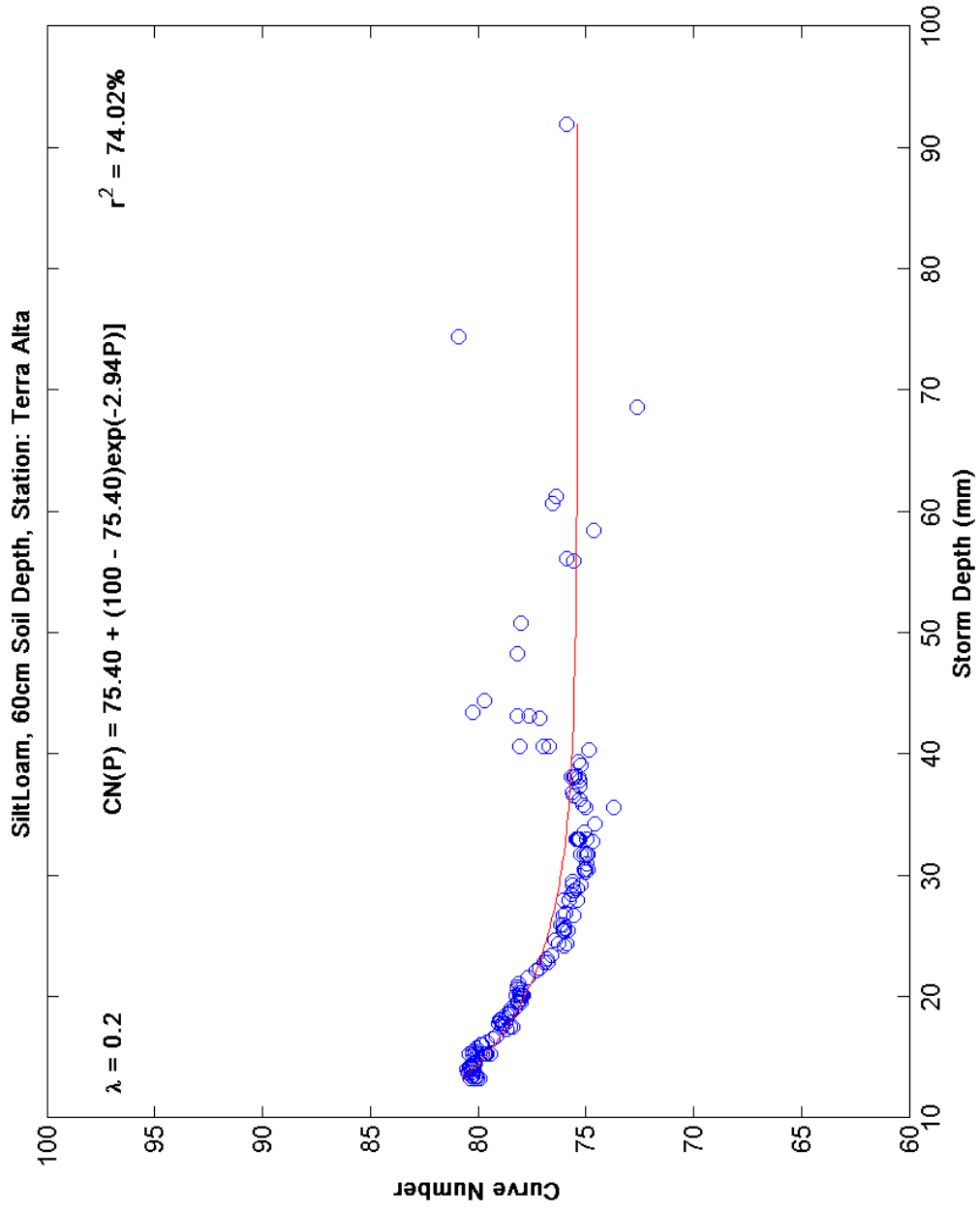


Figure 55. Asymptotic Method, CN vs. Storm Depth, Terra Alta, 60 cm Soil Depth

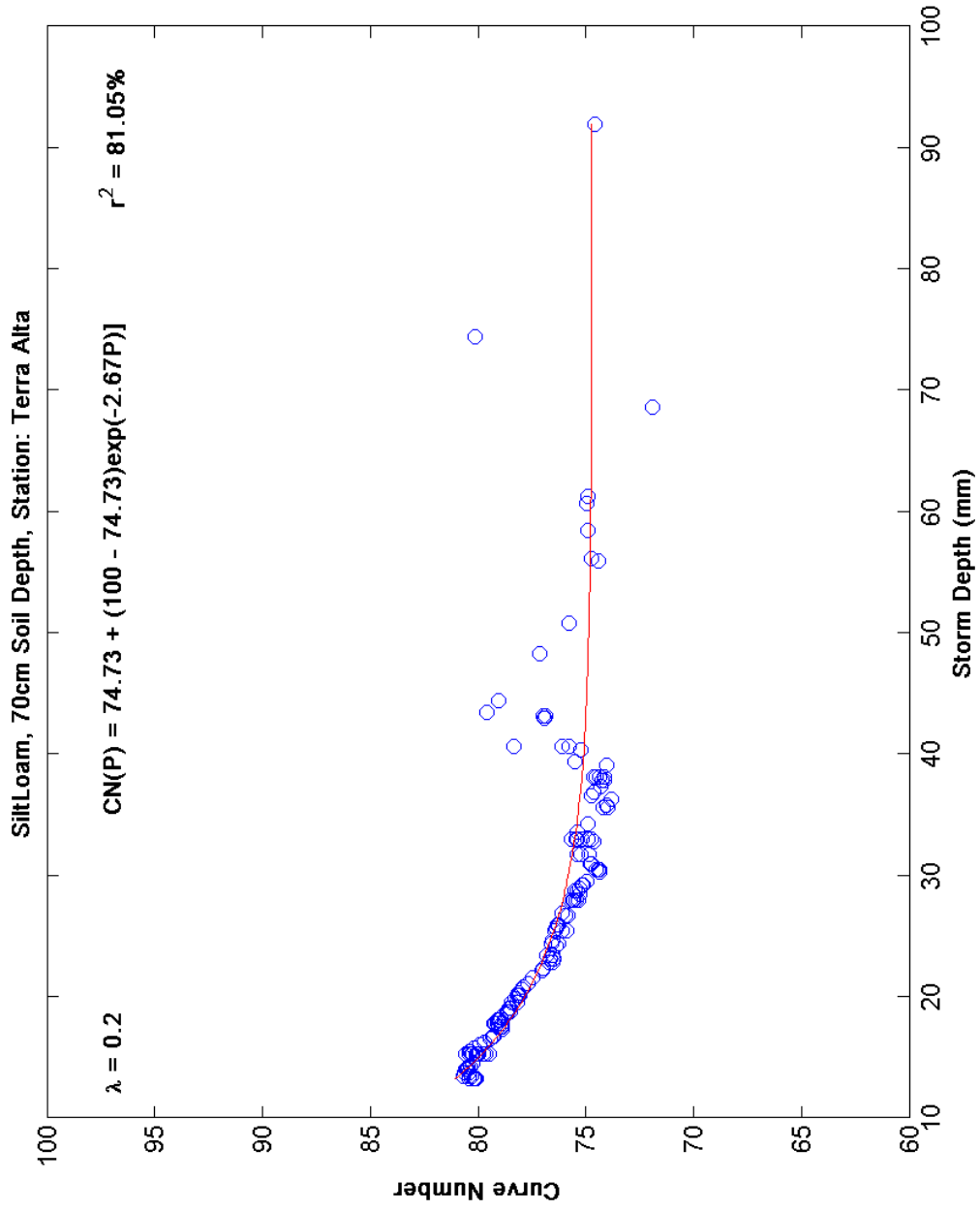


Figure 56. Asymptotic Method, CN vs. Storm Depth, Terra Alta, 70 cm Soil Depth

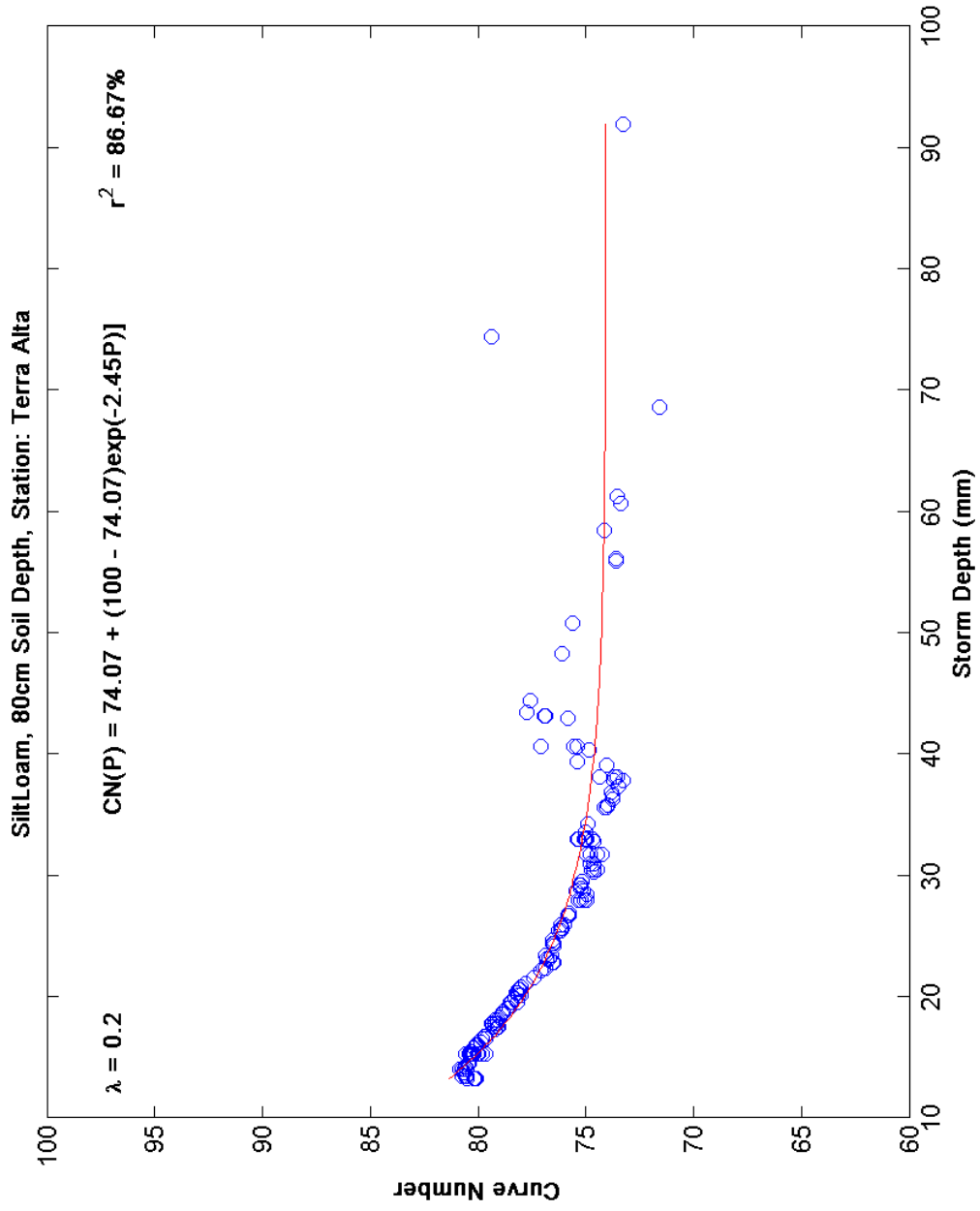


Figure 57. Asymptotic Method, CN vs. Storm Depth, Terra Alta, 80 cm Soil Depth

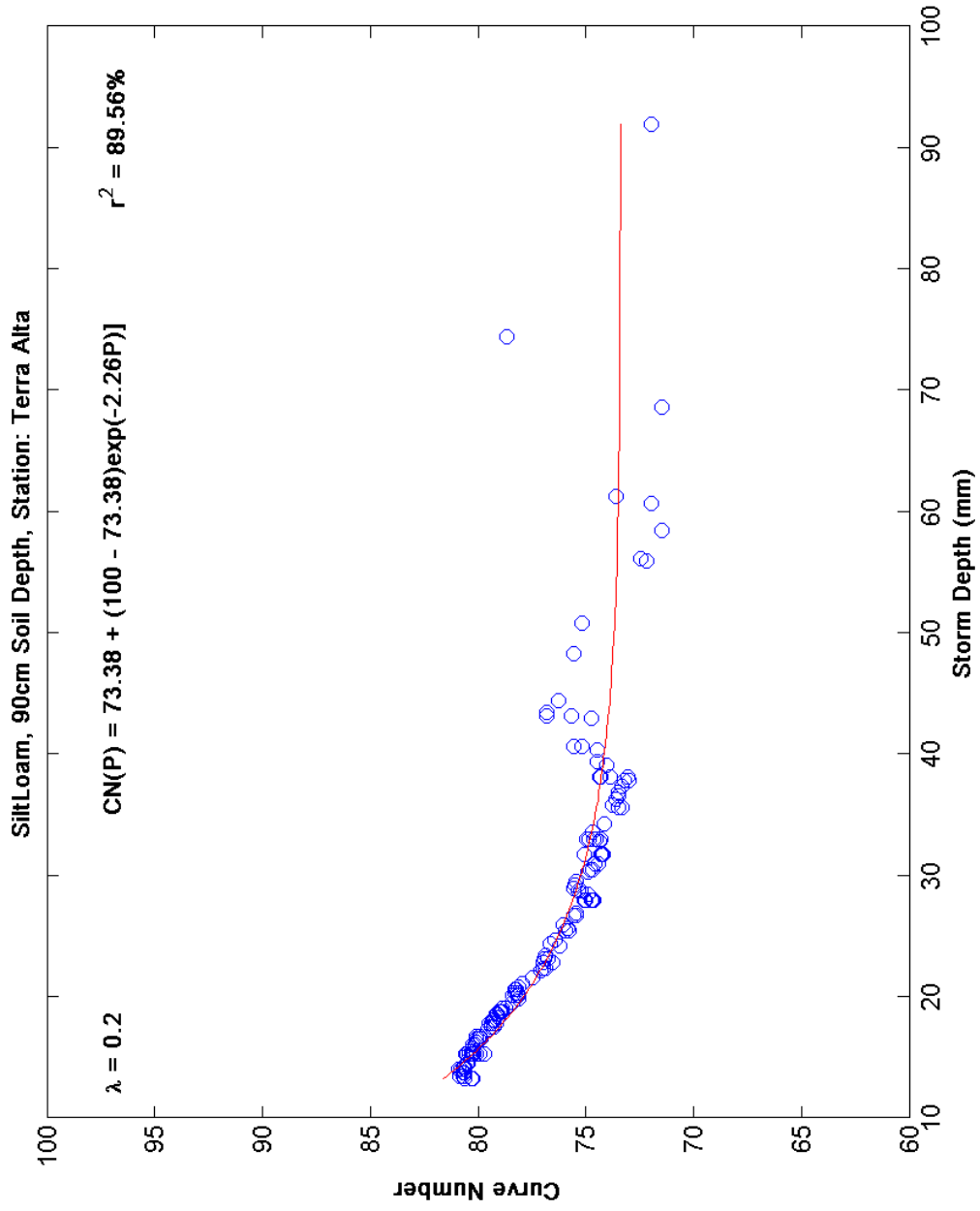


Figure 58. Asymptotic Method, CN vs. Storm Depth, Terra Alta, 90 cm Soil Depth

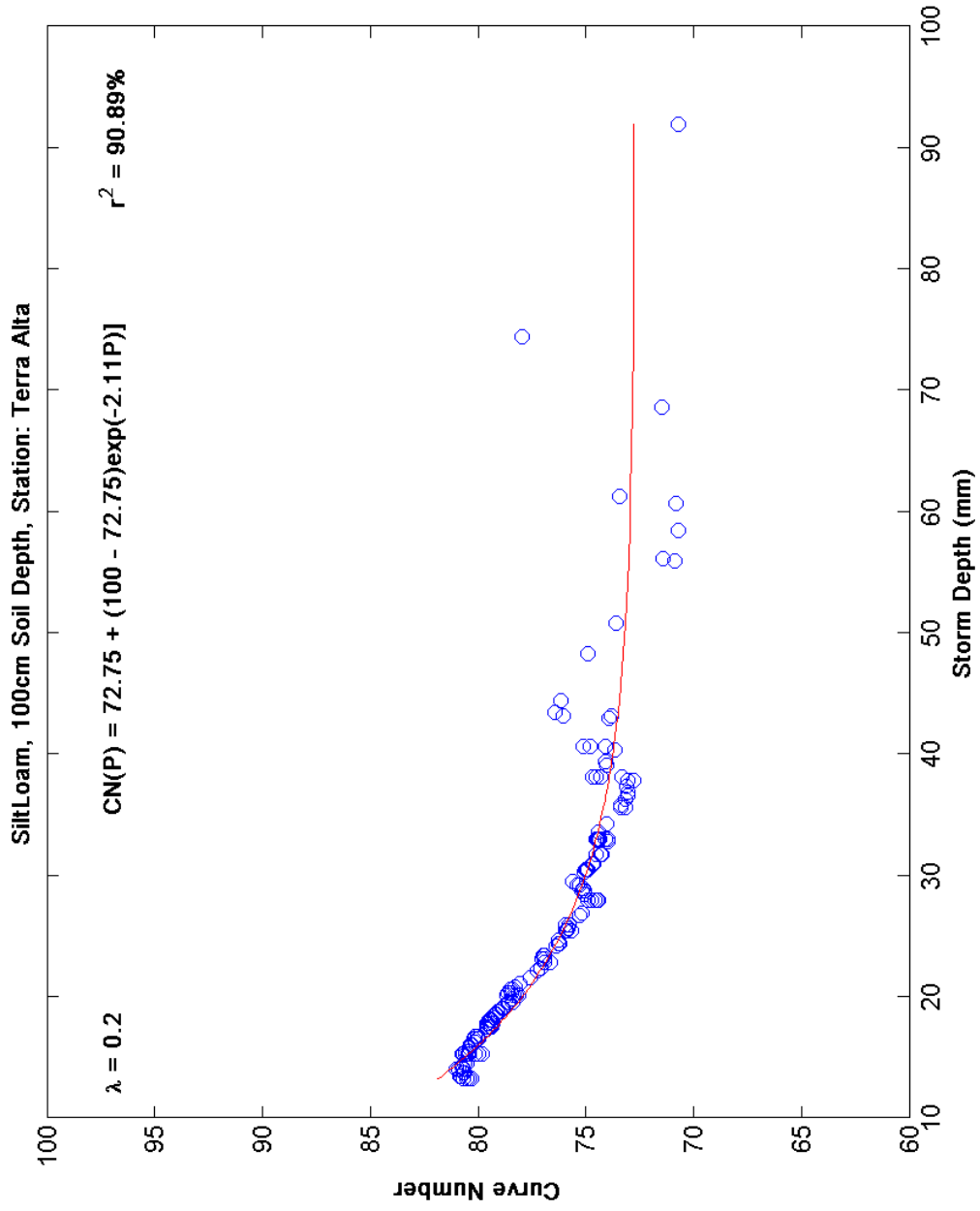


Figure 59. Asymptotic Method, CN vs. Storm Depth, Terra Alta, 100 cm Soil Depth

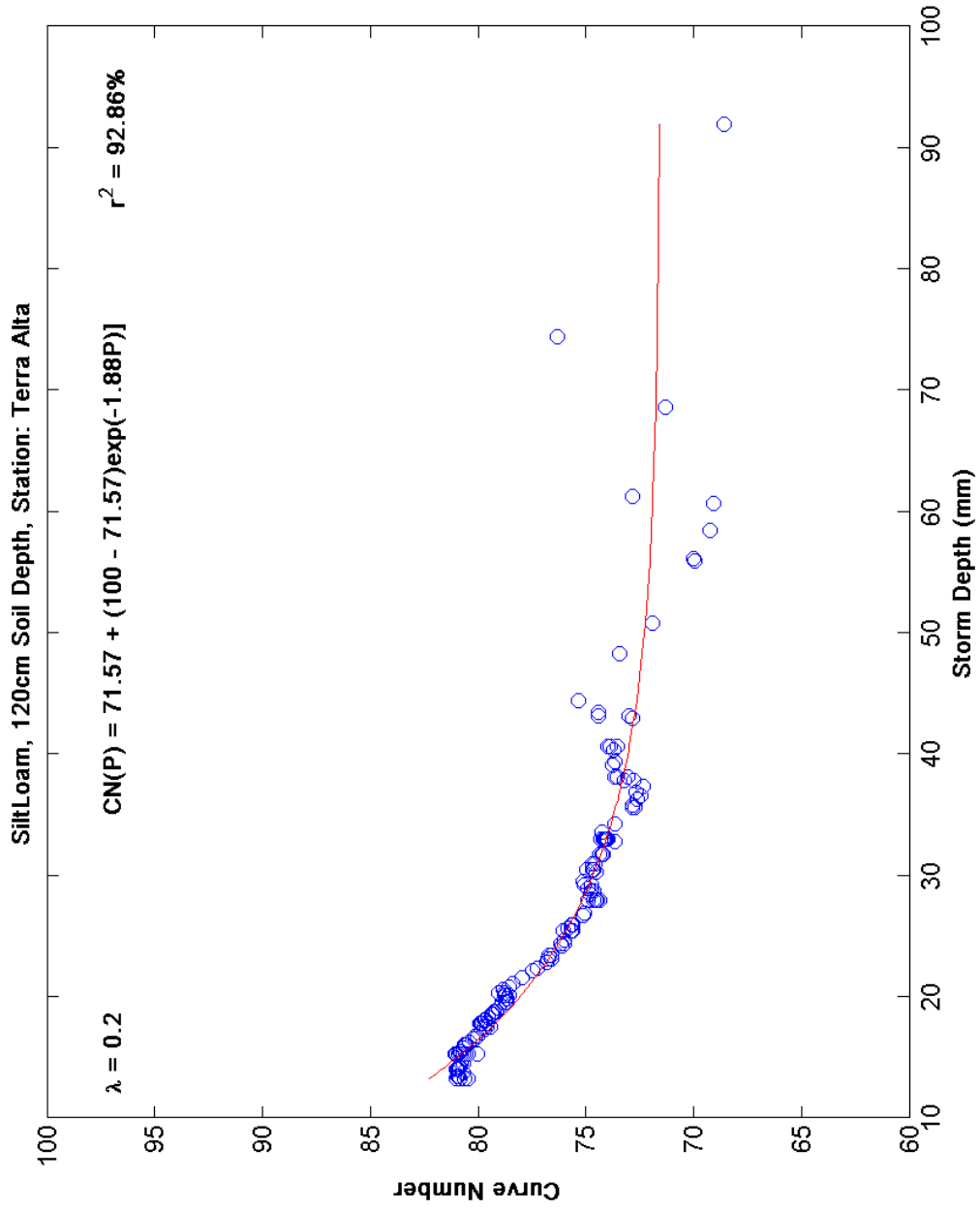


Figure 60. Asymptotic Method, CN vs. Storm Depth, Terra Alta, 120 cm Soil Depth

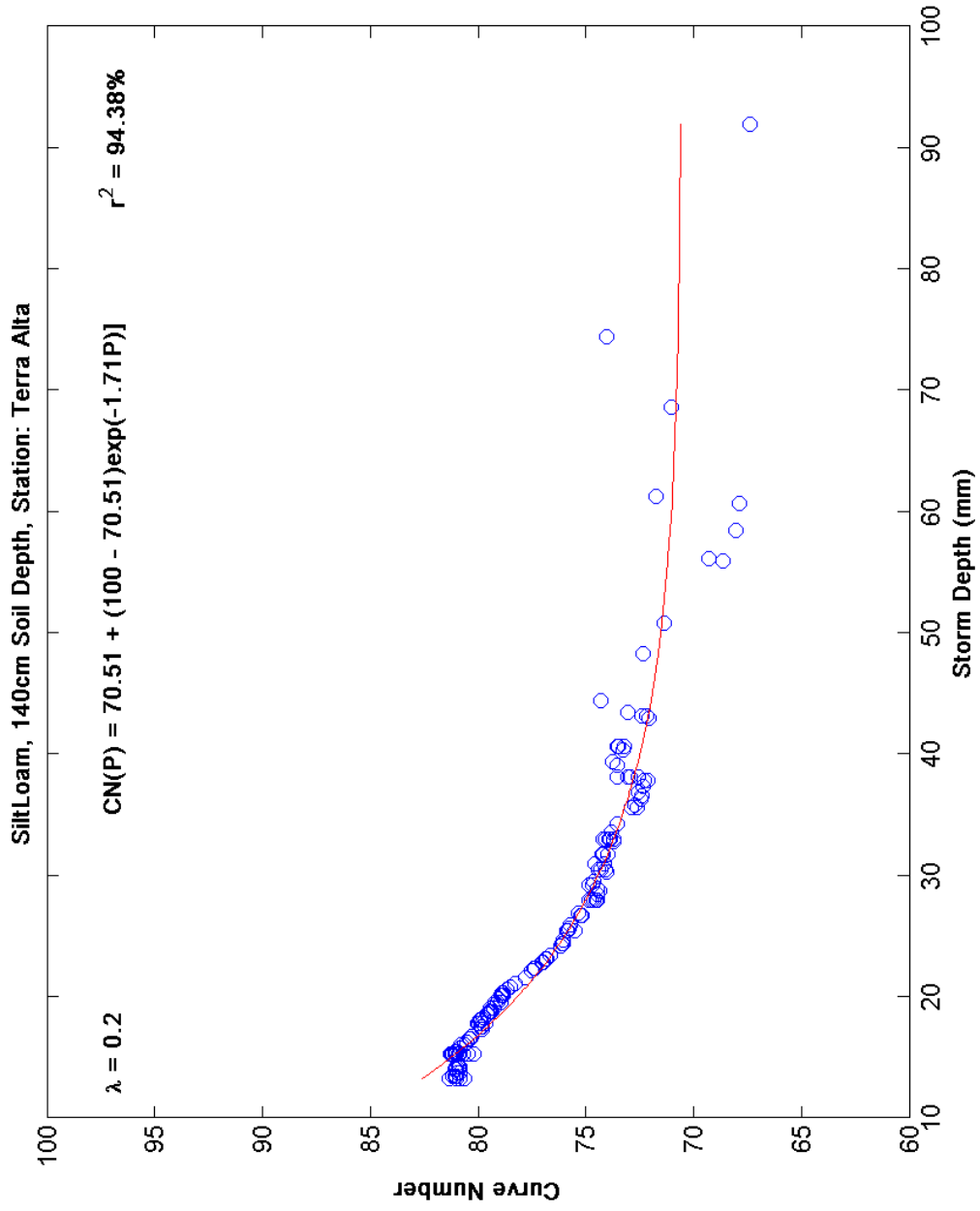


Figure 61. Asymptotic Method, CN vs. Storm Depth, Terra Alta, 140 cm Soil Depth

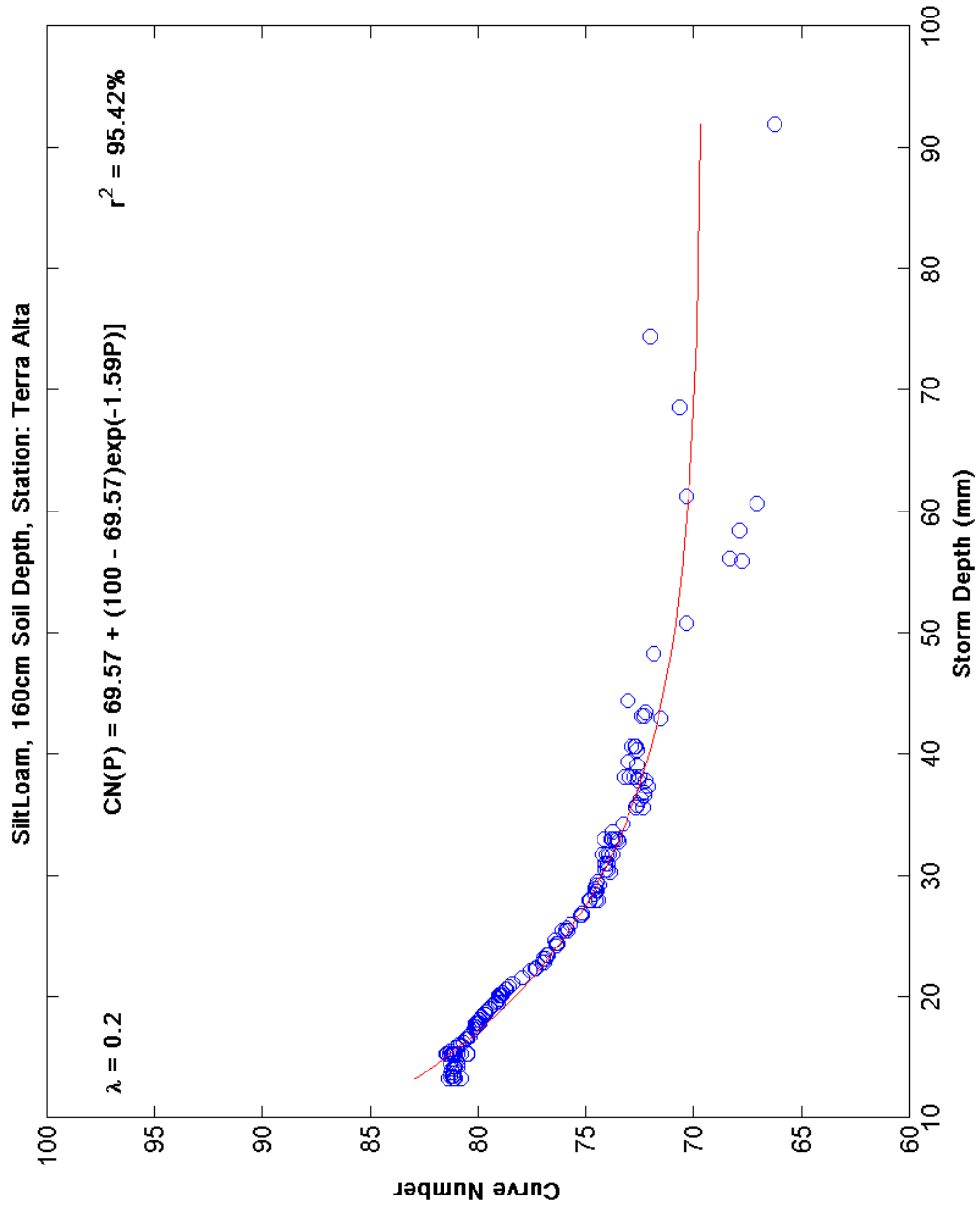


Figure 62. Asymptotic Method, CN vs. Storm Depth, Terra Alta, 160 cm Soil Depth

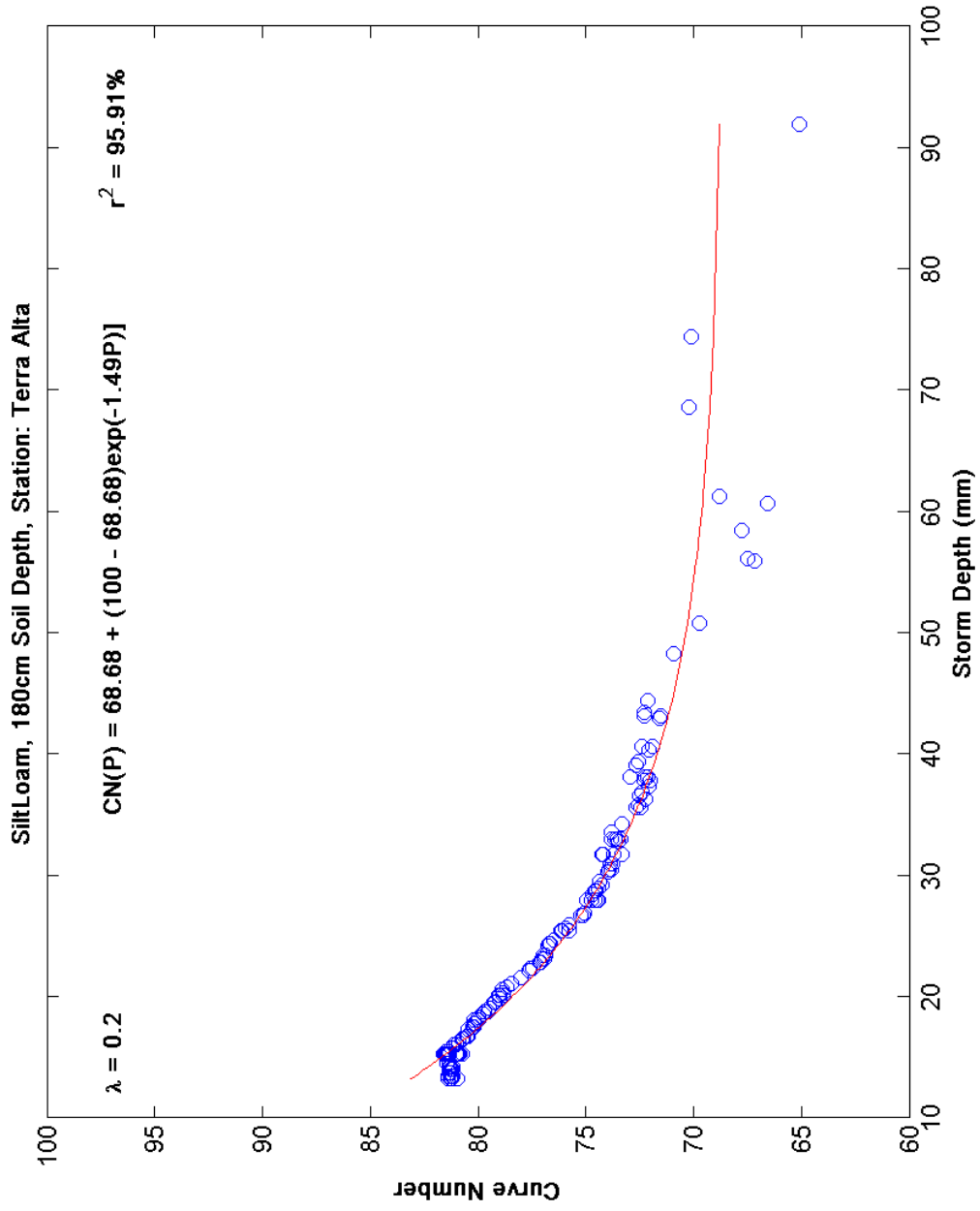


Figure 63. Asymptotic Method, CN vs. Storm Depth, Terra Alta, 180 cm Soil Depth

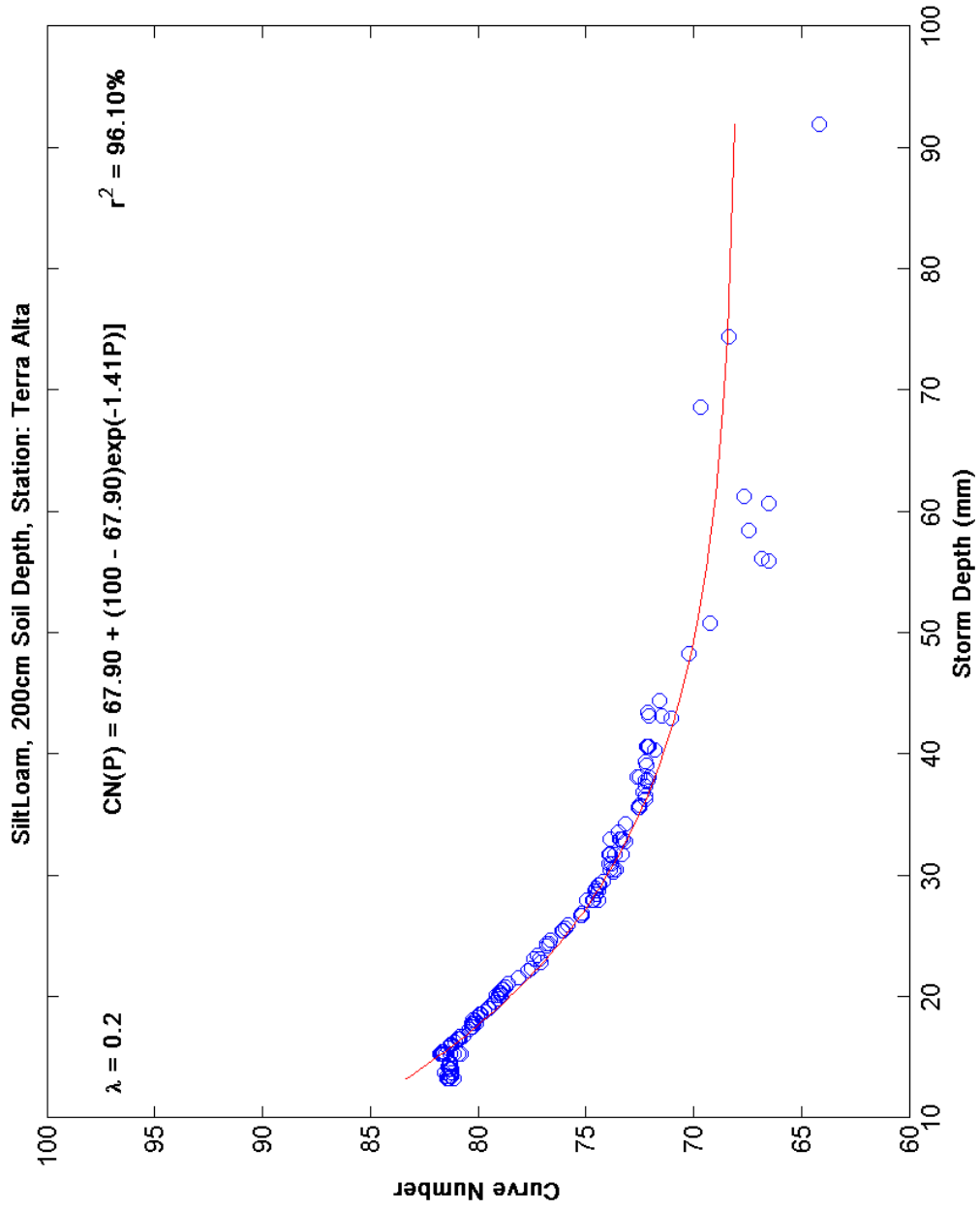


Figure 64. Asymptotic Method, CN vs. Storm Depth, Terra Alta, 200 cm Soil Depth

Table 6 lists the average CN_{∞} values over all soil depths for each gage and soil type. The Dunlow gage value is approximately 2 points higher for the Clay Loam and 3 points higher for the Silt Loam than the other 3 gages (except for the Silt Loam at Terra Alta with $\lambda = 0.05$).

Table 6. Average CN_{∞} Values for Each Gage

	0.2		0.05	
	Clay Loam	Silt Loam	Clay Loam	Silt Loam
Terra Alta	86.98	75.59	83.60	68.34
Beckley	86.35	73.82	82.84	61.51
Elkins	86.85	73.86	82.81	59.83
Dunlow	88.30	76.77	87.27	64.87

This may be explained by noting that on average, according to Figure 11, a greater percentage of the rainfall at the Dunlow gage fell within a shorter time interval (the first five hours) for each storm compared to the other gages. This increased rainfall intensity resulted in slightly lower mean infiltration depths for each soil depth (Figure 65) compared to the other three gages. This effect is analogous to that produced by the WDM Triangular distribution in the Cyclic Method.

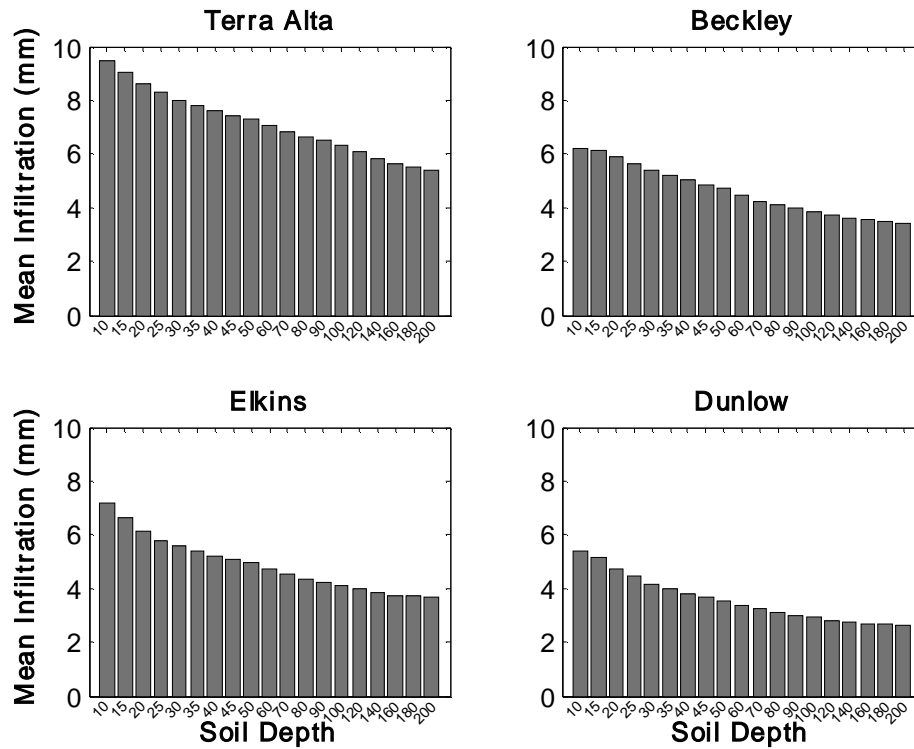


Figure 65. Asymptotic Method Mean Infiltration vs. Soil Depth, Clay Loam

In comparing Figures 45 and 65 it is apparent that in the Cyclic Method, the mean infiltration depth increases with soil depth while the opposite is true for each gage site in the Asymptotic Method. This can be explained by noting the difference between the antecedent lower zone storage (LZS_i) between the methods for each rainfall-runoff event. In the Cyclic Method, LZS_i was calculated using Equation 13 based on soil depth, residual moisture content, and the initial moisture content according to ARC II (Rawls and Brakensiek, 1986). In the Asymptotic Method, LZS_i was determined at the hour preceding each selected storm event. The calculated LZS_i (Cyclic) was consistently lower than the simulated LZS_i (Asymptotic), especially for the greater soil depths. An example comparison is shown in Figure 66 below for the 200 cm soil depth. According to Equation 14, an increase in the ratio of the lower zone storage value (LZS_i) to the lower zone nominal capacity parameter ($LZSN$), results in an increase in the infiltration capacity. Therefore, in the Cyclic Method, the infiltration capacity was substantially

greater than that in the Asymptotic method due to the relatively low value of initial soil moisture storage (LZS_i) predicted by the ARC II condition.

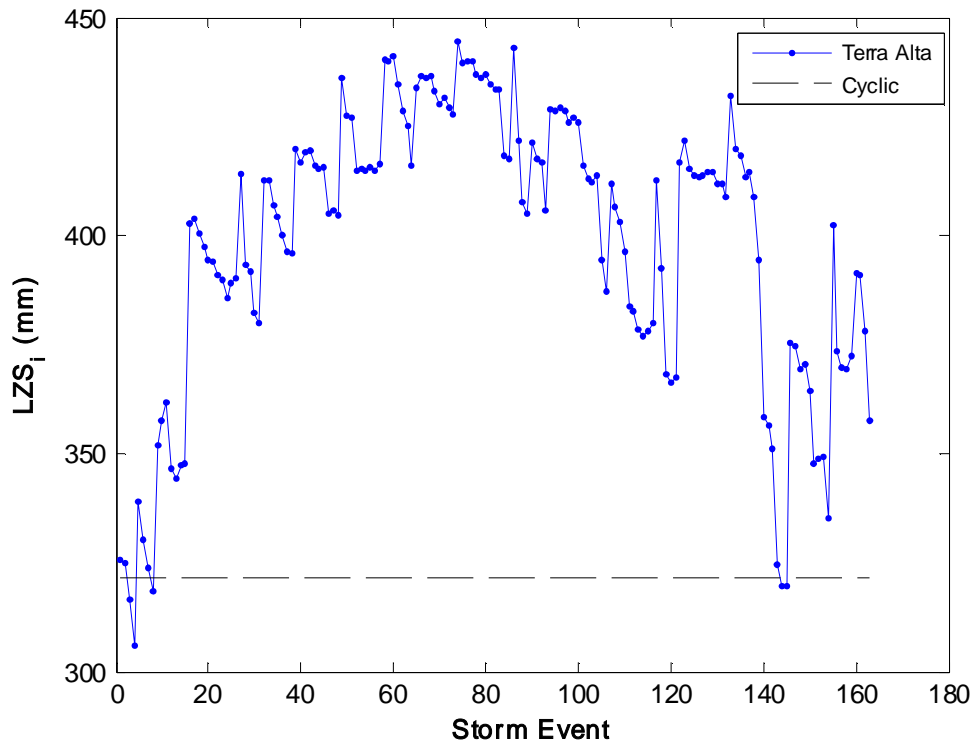


Figure 66. Antecedent LZS_i , Clay Loam, 200 cm Soil Depth, Terra Alta

Finally, in comparing Figures 30-37 to 38-45, it is evident that $\lambda = 0.05$ tends to decrease the CN at low storm depths, as demonstrated by the Cyclic Method. It should also be noted that the Dunlow gage has the shortest record length with the lowest maximum storm depth (47 mm). This was at least 20 mm less than the maximum depths at the other gages.

3.3 Conclusions

By comparing the results of each CN calculation procedure, it is apparent that the CN is dependent on all of the computational variables listed in Table 3. In the Cyclic and Asymptotic methods, the CN's decreased with increasing soil depth due to the increased soil moisture storage capacity. The Clay Loam soil resulted in higher CN's than the Silt Loam because of the lower value of the INFILT parameter governing the infiltration rate

in the Clay Loam. The CN's also vary with storm depth, typically approaching a constant value beyond some threshold depth.

Perhaps the most unanticipated result of this study is the apparent dependence of the CN on storm distribution. In the Cyclic Method, the distributions consisting of the high intensity hourly rainfall intervals (Type II and WDM Triangular) tended to result in greater variation of the CN with storm depth. This effect was also seen in the Asymptotic Method. Figures 7 and 11 demonstrate the similarity in the distributions of the selected storms from all gages to the Type II and WDM Triangular shapes. In each method, the twenty-four hour distribution that allocates the most of rainfall in the shortest time (Dunlow in the Asymptotic Method and WDM Triangular in the Cyclic Method) resulted in the highest CN values and the lowest mean infiltration depths.

These findings suggest that the variability of the CN in time cannot be explained by antecedent soil moisture or rainfall alone and therefore, the use of the CN method in continuous modeling does not appear to be appropriate. It is suggested that a distinction be made between the classic CN and continuous CN's presently in use by using a different symbol such as CN*.

4.0 Summary and Conclusions

The use of the CN as a simplification of several parameters in a comprehensive watershed model (HSPF) was investigated with respect to the analysis of the cumulative hydrologic impacts of surface coal mining in West Virginia. A soil physics model was developed to act as a method of translation between CN's and HSPF parameters based on soil hydraulic properties. Curve Numbers were calculated from theoretical HSPF watersheds (parameter sets) using two numerical methods. The first method is based on the soil physics model and uses cyclic storm input to calculate CN's as a function of soil moisture. The second method is based on Hawkins' Asymptotic method of CN determination (1993) where CN's are calculated from ordered rainfall-runoff pairs. Each method found the CN to be dependent on a number of computational watershed variables including soil type, soil depth, storm depth, and storm distribution. The effect of the initial abstraction ratio of 0.05 vs. 0.2 was found to reduce the bias of high CN values at low storm depths.

These findings suggest that the hydrologic information inherent to the CN method is insufficient for the CN to adequately represent multiple HSPF parameters. Because of its apparent dependence on several watershed variables which are naturally irregular in space and time, the CN method appears to be unsuitable for continuous rainfall-runoff predictions. Application of the CN method should be limited to single, event-based runoff estimation as described in the original development of the method. The development of a translation methodology between the CN and HSPF parameters based on soil physics was successful, however, the use of the CN method with HSPF to simulate the hydrologic impacts of mine sites is not recommended. The effects of long-term land use change in general are best quantified by gathering actual rainfall-runoff data. Accurate simulations of the effects of surface mining using HSPF (or other continuous models) in West Virginia will require several years of hydrologic and meteorological time series records from the mine sites themselves. These records would be extremely valuable for studying the cumulative hydrologic impacts of coal mining by providing the ability to calibrate continuous models to observed data.

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Appendix A. HSPF Calibration, Verification, and Parameter Optimization Study

This is an unpublished technical report prepared by, Dr. Robert N. Eli¹, Samuel J. Lamont¹, and Elena Hoeg². Funding for this research was provided by the West Virginia Department of Environmental Protection. Additional support was provided by the Office of Surface Mining, Jim Sams of the USGS, Pittsburgh, PA, and Kate Flynn of the USGS, Reston, Virginia.

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Review of HSPF Background and Related Applications

HSPF was selected as the hydrologic model for CHIA of mine-impacted watersheds in the state of West Virginia because of its wide use and acceptance as a joint watershed and stream water quality model. It is a comprehensive, continuous watershed simulation model, designed to simulate all the water quantity and water quality processes that occur in a watershed (Bicknell, et al., 2001). This includes sediment transport and movement of contaminants overland and through the stream channel system. HSPF has its origins in the Stanford Watershed Model (SWM) developed by Crawford and Linsley (www.hydrocomp.com). This latter model was the first truly comprehensive land surface and subsurface hydrologic processes model that treated every component of the hydrologic cycle. It has been widely adopted in various forms and its hydrologic components have been included in related models, such as the Kentucky Watershed Model. Crawford and Linsley further developed the original SWM model and created HSP, the Hydrocomp Simulation Program, which included sediment transport and water quality simulation. Hydrocomp also developed the ARM (Agricultural Runoff Management Model) and the NPS (Nonpoint Source Pollutant Loading Model) for the EPA (U.S. Environmental Protection Agency) during the early 1970's. In 1976, EPA commissioned Hydrocomp, Inc. to develop a set of simulation modules in standard Fortran that would handle all the functions handled by HSP, plus those within two additional models, ARM and NPS. The intention was to produce a modeling system that was easy to maintain and modify. The result was HSPF, which can be applied to most watersheds using commonly available meteorologic and hydrologic data. HSPF has been applied to a variety of watershed studies, including the U.S.EPA Chesapeake Bay Program, Carson - Truckee River (California, Nevada), Minnesota River Assessment Project, Florida Water Management District, King Co. Washington Management Plan, and others (Donigian, 2003). Other work that relates specifically to various aspects of the calibration methodology used here includes Sams and Witt (1995), and Dinicola (2001). Sams and Witt (1995) utilized HSPF to model two surface-mined watersheds in Fayette County, Pa. The significance of this latter study is the location of these two watersheds, located within and just to the north of the Big Sandy calibration watershed which is one of the calibration watersheds. The Stony Fork Basin is a sub-basin of Big Sandy, and the Poplar Run Basin is located just 15 miles to the north of Big Sandy. The geology, soils,

topography, and land cover of these two watersheds are very similar to the characteristics of many of the trend station, calibration, and verification watersheds used in the CHIA project. Therefore, the fitting parameters as determined by Sams and Witt (1995), were adopted as a starting point in the calibration processes for the CHIA project. Additional studies of note are those by Al-Abed, et al., (2002), Lohani, et al., (2002), Martin, et al., (1990), Riberio (1996), and Srinivasan, et al., (1998).

Summary of HSPF Basic Capabilities and Characteristics

The HSPF model has the following general characteristics:

- It is a continuous simulation model (It can simulate streamflow for many years at hourly time increments).
- It can be applied to natural or developed watersheds (including those with surface and underground mine sites).
- Model components simulate both the land surface and subsurface hydrology and water quality processes.
- HSPF utility programs provide time series data management, statistical analysis tools, and graphic display of results.
- Both stream and lake hydraulics and water quality processes can be simulated.
- HSPF is the core watershed model in EPA BASINS and the U.S. Army Corps of Engineers WMS modeling system.
- Development and maintenance of HSPF related software is sponsored by EPA and USGS.

There are three application modules that make up the core of the HSPF hydrologic model (each also includes several sub-modules of importance):

- 1) PERLND (Simulate a Pervious Land segment)
 - a) ATEMP (Correct air temperature for elevation difference)
 - b) SNOW (Simulate the accumulation and melting of snow and ice)
 - c) PWATER (Simulate water budget for pervious land segments)
- 2) IMPLND (Perform computations on a segment of impervious land)
 - a) ATEMP (Same as in PERLND above)
 - b) SNOW (Same as in IMPLND above)
 - c) IWATER (Simulate water budget for impervious land segment)
- 3) RCHRES (Perform computations for a stream reach or mixed reservoir)
 - a) HYDR (Simulate hydraulic behavior)

b) ACIDpH (Simulate mine acid drainage in-stream chemistry)

Of the three application modules above, PERLND and RCHRES were used in the calibration phase of the CHIA project. The PERLND module simulates the watershed areas, with each land cover/land use classification category being described by its own unique set of PERLND parameters. The RCHRES module is applied to each stream reach, which is equivalent to a stream segment in the stream drainage network within a given watershed. Each stream reach has its own unique descriptive parameters, which are applied in the RCHRES module. The IMPLND module is for the purpose of simulating impervious areas, such as urban areas. This module was not used since no urban areas larger than a few percent of the total watershed area are encountered in the CHIA project.

CHIA Calibration and Verification Watersheds

Watershed Selection

The hydrologic component of the project involves the fitting of HSPF to each of the 235 Trend Station Watersheds identified by WVDEP. They have boundaries defined by stream water quality sampling points, or Trend Stations, located at the watershed outlets. These stream water sampling points generally do not coincide with USGS stream gaging locations that are required for the model calibration process. Therefore, model calibration must be conducted using watersheds that have a gaging station at their outlet, and are also representative of the hydrologic characteristics found in CHIA watersheds. An additional factor is the obvious impracticality of individual calibration of 235 watersheds, regardless of gaging data availability. The only practical approach to finding a set of model parameters for each of the 235 trend station watersheds is to calibrate the model to a selected few watersheds that contain representative characteristics of the whole population of watersheds. It is then assumed that watersheds with similar characteristics have similar model parameters representing those characteristics. It is therefore possible to calibrate a limited number of watersheds as long as their hydrologic characteristics are simulated as separable components in the hydrologic model. The suitability of the parameter sets determined during calibration is tested using a set of verification watersheds that are also representative of the CHIA watersheds. This calibration strategy follows that recommended by Donigian (2002), and successfully employed by Dinicola (1990, 2001). The Dinicola (2001) study involved 12 small watersheds in King and Snohomish Counties, in and near Seattle, Washington. The purpose of this latter study was to model the effects of urbanization on watershed response. Five of the watersheds were selected for use in calibration, characterized by various degrees of development. The calibration process proceeded with the intent to arrive at a consistent set of parameters across all 5 watersheds for each land use category. The study was successful in that it demonstrated that satisfactory model performance could be achieved by using common land use categories with single valued parameter sets. The approach used in the CHIA calibration study follows Dinicola's lead in maintaining a single valued set of model parameter values for each land use category.

The calibration and verification watersheds lie within the coal regions and either encompass or are adjacent to trend station watersheds. Figure 1 shows the locations of

the trend station watersheds within the state of West Virginia, including the five watersheds selected for calibration purposes. It will be noted that the Twelve Pole Creek, Clear Fork, Buffalo Creek, and Big Sandy watersheds contain trend station watersheds in whole or in part. Big Sandy lies partially in the state of Pennsylvania, and therefore only the West Virginia portion contains trend station watersheds. Tygart Valley at Elkins does not contain trend station watersheds, but lies adjacent to trend station watersheds on its western boundary. Figure 2 shows the location of five verification watersheds which are used to test the modeling parameters determined in the calibration process. These include Big Sandy (same as the calibration watershed, except using a different meteorological record), Tygart Valley at Belington, Tygart Valley at Daily, Piney Creek, and Panther Creek. It will be noted that the two Tygart Valley verification watersheds are a superset and subset of Tygart Valley at Elkins, respectively. These latter two verification watersheds are defined by different gaging locations along the same stream, and hence share a portion of the same watershed. The Big Sandy watershed is present in both the calibration and verification watershed groups to provide for error checking.

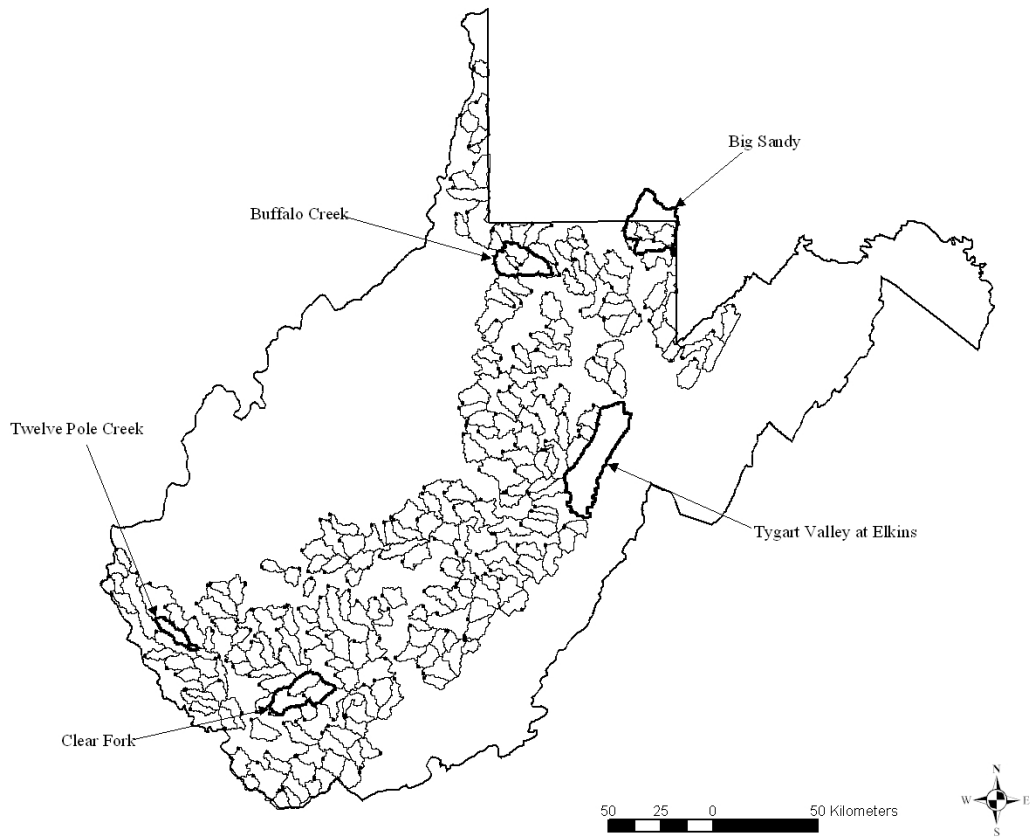


Figure 1 : West Virginia CHIA Trend Stations and Calibration Watersheds.

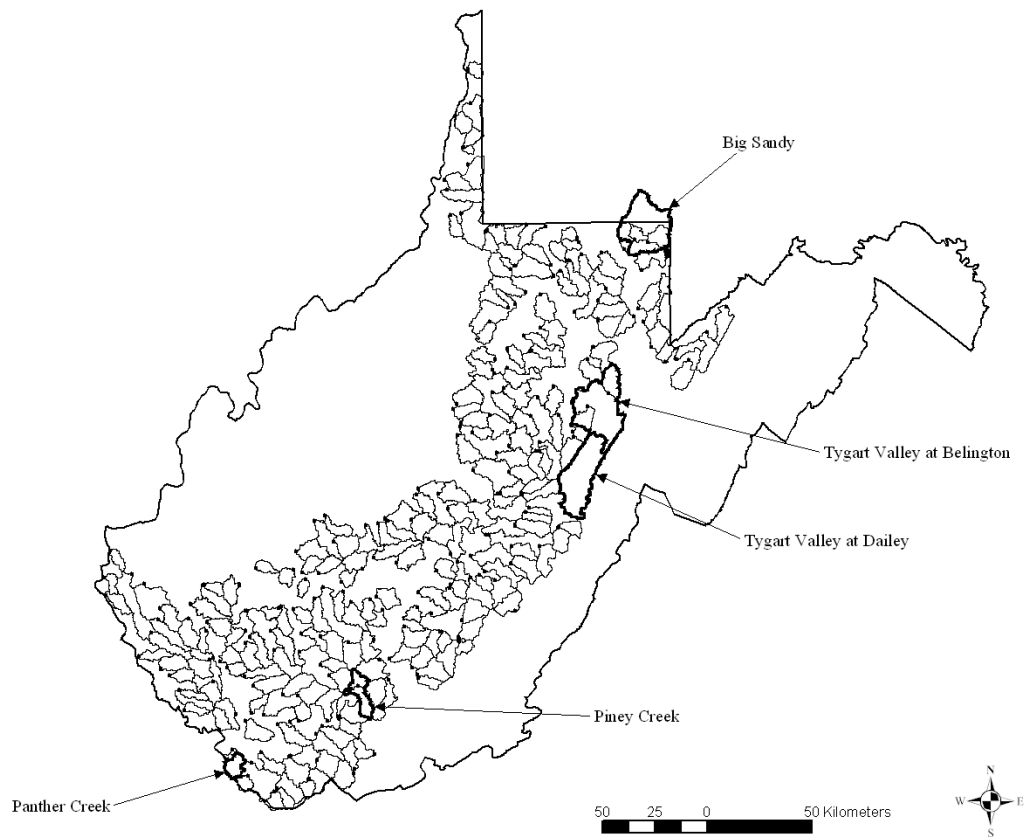


Figure 2 : West Virginia CHIA Trend Stations and Verification Watersheds

Watershed Characteristics

The calibration and verification watersheds, shown in Figures 1 and 2, required stream flow gaging data to support the HSPF model fitting process. Table 1 lists the watersheds along with the available USGS stream flow record and corresponding gage number.

Table 1 : List of Calibration and Verification Watershed Available Gaging Records.

	Watersheds	Stream Flow Record		Gage Number
	Calibration	From	To	
1	Twelve Pole Creek	10/01/1964	09/30/2000	03206600
2	Buffalo Creek	06/03/1907	09/30/2000	03061500
3	Tygart River at Elkins	10/01/1944	09/30/2000	03050500
4	Clear Fork	06/28/1974	9/30/200	03202750
5	Big Sandy	05/07/1909	09/30/2000	03070500
	Verification			
1	Panther Creek	08/01/1946	09/30/1986	03213500
2	Tygart River at Belington	06/05/1907	09/30/2000	03051000
3	Tygart River at Dailey	04/20/1915	09/30/2000	03050000
4	Piney Creek	08/21/1951	09/30/1982	03185000
5	Big Sandy	see above		

The land use/cover classifications are based on 1993 GAP data. The classifications used are:

1. Forest
 - a. Steep Slope
 - b. Moderate Slope
 - c. Mild Slope
2. Barren
3. Mined
4. Pasture/Grassland
5. Row Crop/Agriculture
6. Shrubland
7. Surface Water
8. Urban/Developed
9. Wetland

It should be noted that a total of 11 classifications result due to the forested slope sub-categories, which are treated as separate classifications. Table 2 lists the total watershed area and distribution of areas in the forest slope classifications for each of the calibration watersheds, illustrating the predominance of the forest category.

Table 2 : Slope Distribution for Calibration Watersheds

Watershed	Total Area (acres)	Total Forested Area (acres)	% Forested	% Mild Forest	% Moderate Forest	% Steep Forest
Twelve Pole Creek	23646	20402	86	10	16	74
Buffalo Creek	72257	57590	80	19	28	53
Tygart Valley at Elkins	172642	137950	80	16	22	62
Clear Fork	79862	71455	89	7	10	83
Big Sandy	123027	96713	79	61	29	10

Each of the calibration watersheds has a mining history. Figure 3 shows the relative cumulative percentages of surface and underground mining in the calibration watersheds, as documented in annual mine permit application records. It should be noted that significant historical mining is not documented on these watersheds, but is known to be present. As an example, it is known that much of the Pittsburgh coal is mined out under Buffalo Creek watershed, yet it does not appear in the available mapping database. The only mined areas used in the calibration study are those classified as “mined” land in the 1993 GAP database. These latter areas are known not to be accurately classified; however, they were used since there was little to be gained by trying to include other sources of mined land data. Much of the historical mined land is reclaimed, or is overgrown with vegetation, and therefore is now classified as forest, shrubland, or pasture/grassland. Since the purpose of conducting a baseline calibration is to provide a reference condition, against which the effects of new mining can be compared, it serves no useful purpose to try to identify the historical mined areas in an effort to correct the GAP data. The difficulties in trying to treat historical mined areas as a unique classification is not warranted since the HSPF model is a lumped parameter model for which small differentiation in parameters over limited areas has no significant effect on the baseline model output (this behavior was adequately demonstrated during the calibration study to be discussed later). Therefore, the mined classification in the 1993 GAP is retained since it is an integral part of the data set, and its area must be conserved. Likewise, the surface water classification was modeled as a land surface category in order to conserve watershed drainage area. The surface water areas involved are very small and have no impact on the baseline calibration.

Mining Activity for Calibration Watersheds

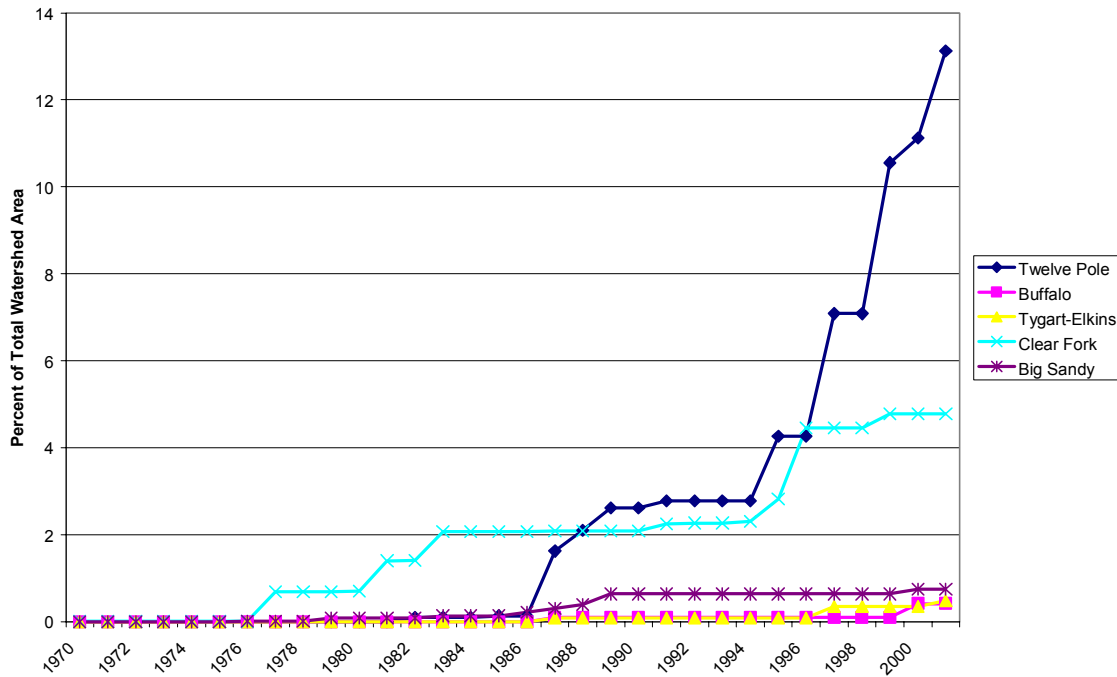


Figure 3 : Calibration Watershed Mining History (from WVDEP Permit Records).

HSPF Meteorologic Data Input Requirements

Meteorologic data required to run HSPF for the calibration process included PET, TEMP, and PREC (potential evapotranspiration, average air temperature, and precipitation). The values for PET and TEMP are estimated from daily maximum and minimum air temperatures (TMAX and TMIN). These data are supplied by NCDC (National Climatic Data Center) and downloaded from the internet (or obtained from a secondary supplier). PET is estimated using a HSPF data utility program called WDMUtil (using the Hamon formula). HSPF uses an hourly time increment for precipitation data input. The precipitation data was supplied under contract by Zedx Inc., which is formatted into average hourly values for each of 5 km grid squares covering the state of West Virginia and portions of surrounding states for the period from 1948 through 2000 (see Figure 4). The daily streamflow data was downloaded from a USGS internet web site. Snow cover was simulated using the temperature-index method option within HSPF.

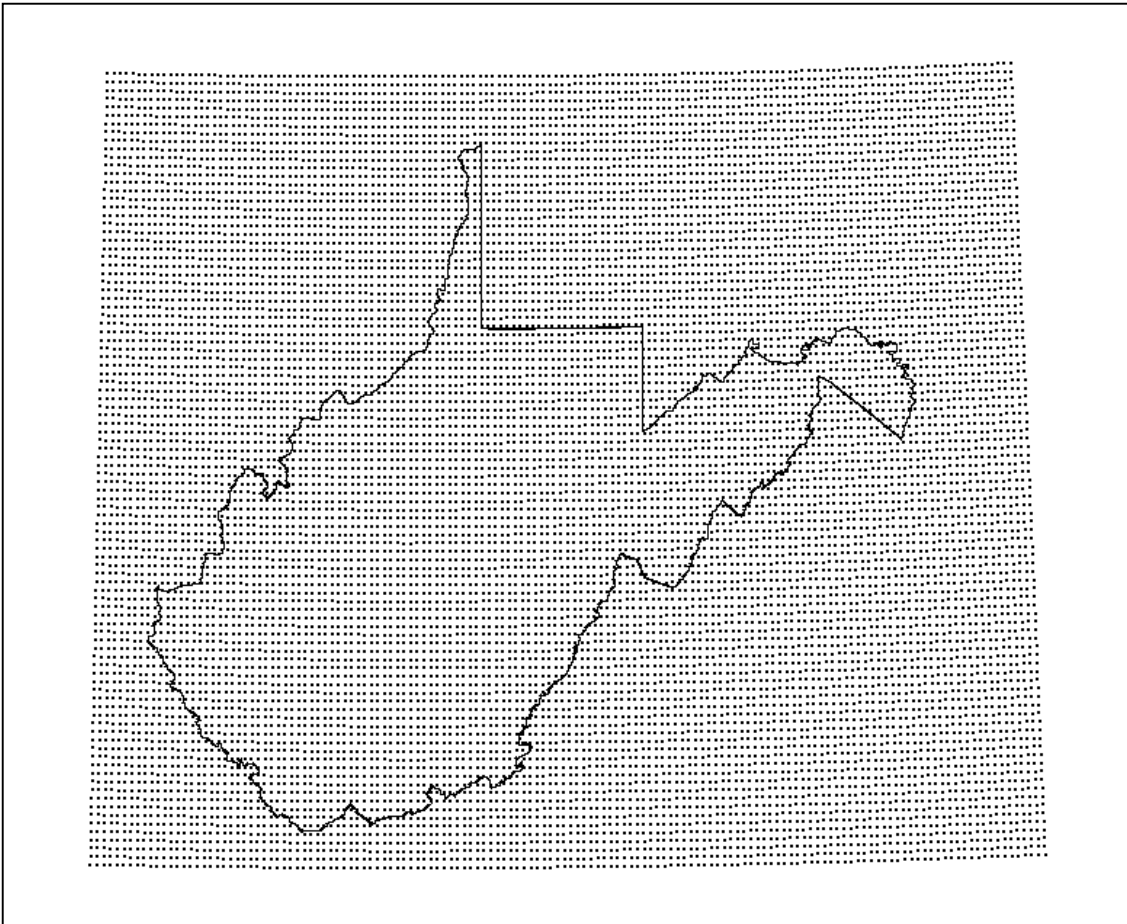


Figure 4 : Geo-located 5 km Grid Square Centers for the Zedx Hourly Precipitation Data.

HSPF Model Calibration and Verification Procedures

Application of EPA BASINS in the Calibration Process

The HSPF model is typically applied to a watershed using BASINS (USEPA, 1999) because of its built-in spatial data base and analysis tools that greatly simplify the input data preprocessing. BASINS automates much of what was formally a very tedious text editing process of building the HSPF user control input (uci) file, by taking the user through a much simpler Windows-based data entry process. The BASINS version of HSPF works reasonably well for general purpose water quality applications but does not have an acceptable acid mine drainage (AMD) water quality (chemistry) modeling capability. The BASINS user interface still requires considerable investment in user time to overcome a steep learning curve. It requires familiarity with four separate pieces of software to prepare the input data, edit the user control input (uci) file, then execute the model, and finally, analyze the results. These latter shortcomings has been addressed by expanding the capability of WCMS to include all of the HSPF modeling and data analysis tools in a single simplified user interface.

It was necessary to conduct the trend station watershed calibration study using BASINS to process the spatial data, and to generate the uci (user control input) files, since the corresponding WCMS tools were still under development during the initial

phases of the CHIA project. In its default form, BASINS provides for automated watershed closure and subdivision using the 1:100,000 scale national DEM. Initially, corrected 1:24,000 DEM (30 m resolution) coverage for West Virginia was substituted to provide the resolution thought needed for the WVDEP-CHIA HSPF model.

Additionally, the existing DLG of the stream networks within BASINS was upgraded to the 14 digit NHD (National Hydrologic Database standard). These modifications then matched the topographic and stream network data resolution to that of the standard 7.5 min. USGS quadrangle map, instead of the 1:100,000 scale map base. Ironically, limitations within the HSPF code ultimately dictated a return to a 1:100,000 scale, and a corresponding 8 digit NHD stream network resolution. As will be presented later, modeling accuracy was not significantly affected due to the lumped parameter characteristics of HSPF.

Watershed Segmentation for HSPF Model Calibration

Segmentation of each calibration watershed into sub-watersheds was based on selection of a sub-watershed size that yields a maximum of approximately 10 sub-watersheds. This was a requirement for calibration only, since the calibration method used limits the number of sub-watersheds and their associated stream segments. Figure 5 shows the Twelve Pole Creek watershed segmented based on a 100 hectare sub-watershed area threshold, yielding 59 sub-watersheds. This is approximately equivalent to the resolution initially planned for use in the WCMS-HSPF model implementation. This is compared to the segmentation of Twelve Pole Creek using a 600 hectare threshold area, as shown in Figure 6, which is representative of the approximate number of sub-watersheds used for the 5 calibration watersheds. Experience of other investigators (personal communication, Kate Flynn, USGS, 2003), points out that the model calibration parameters are not significantly different for coarse segmentation as compared to a fine (high resolution) segmentation of the watershed, as long as the grouped option of assigning the PERLND properties is used (explained later). Independent testing of this thesis was confirmed by simulation comparisons. Figure 7 shows the output of a HSPF simulation for Twelve Pole Creek using 59 and 5 sub-watersheds, respectively, with all other parameters and inputs held constant. The only noticeable difference between the hydrographs is the slightly higher estimation of storm peaks by the 5 sub-watershed model, which is considered of minor significance for calibration purposes. The calibration and verification HSPF watershed models used the 600 hectare threshold criteria for segmentation in order to meet the requirements of the HSPEXP software used for the calibration process (Users Manual, HSPEXP, (1994)). Final segmentation of the trend station watersheds will be done at the 1:100,000 map scale and 8 digit NHD stream network resolutions. This level of detail corresponds to that necessary to sufficiently represent the watershed hydrology and to support the modeling of in-stream chemistry of mine acid drainage.

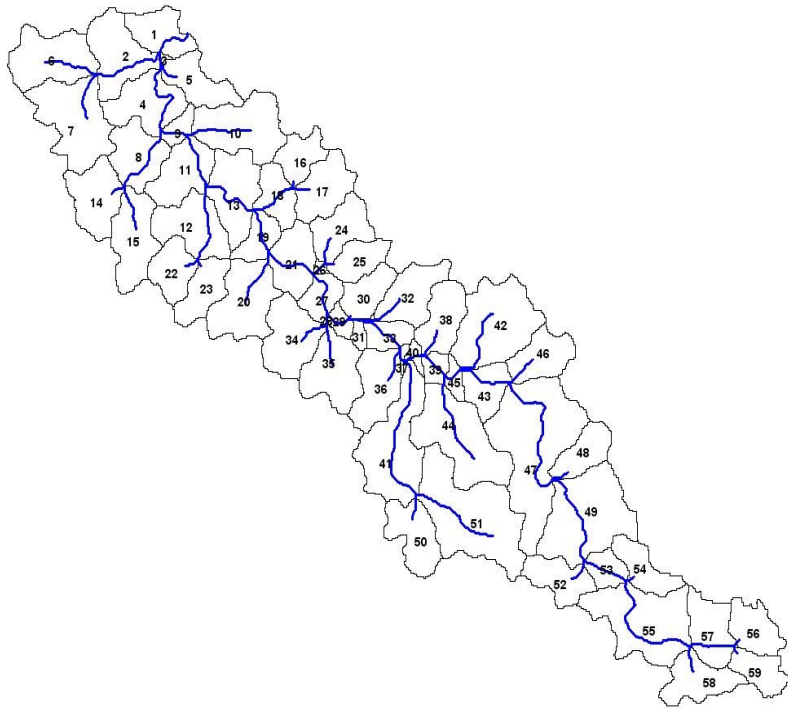


Figure 5 : Twelve Pole Creek Watershed with a 100 ha Threshold Area (59 Sub-Watersheds)



Figure 6 : Twelve Pole Creek watershed with a 600 ha threshold area (5 sub-watersheds)

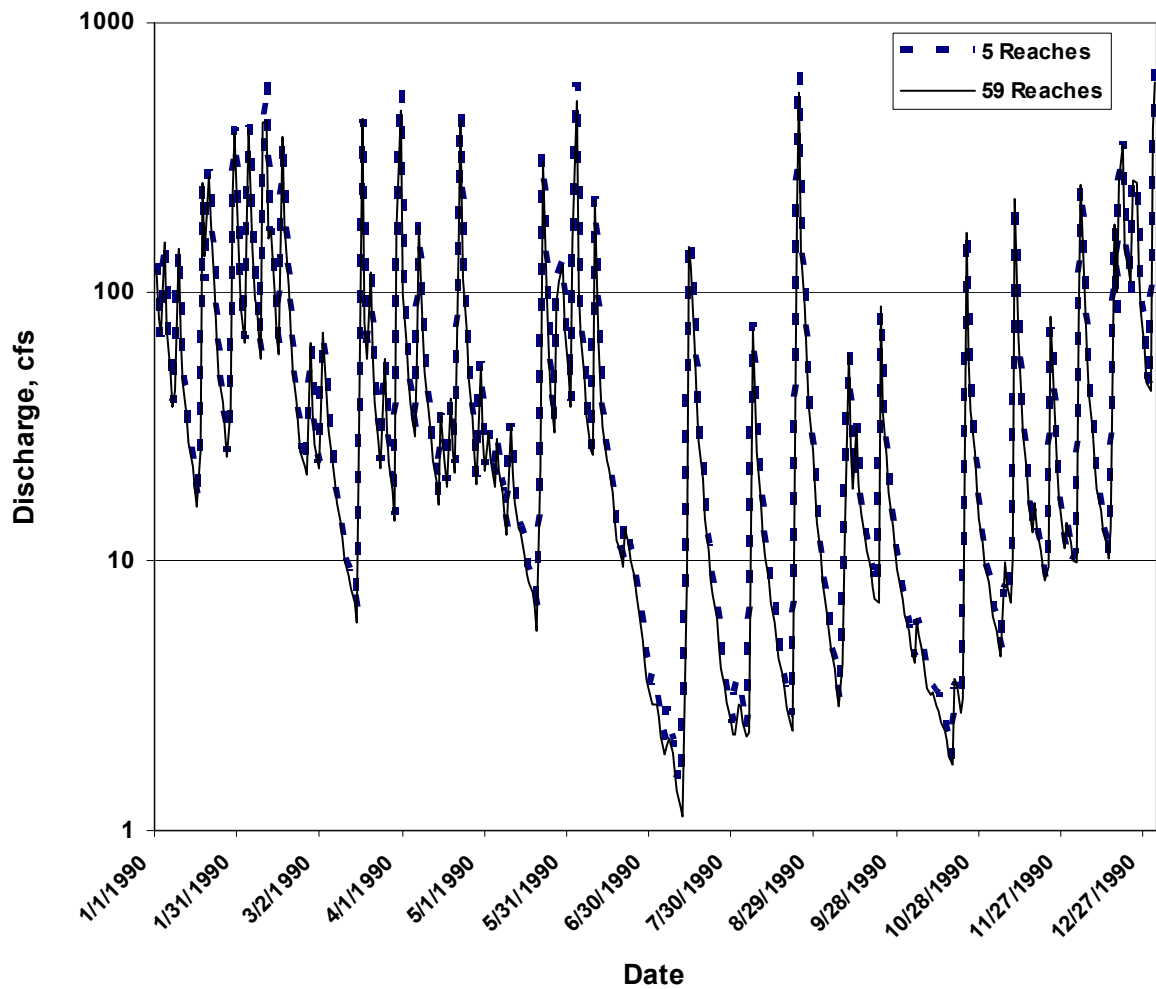
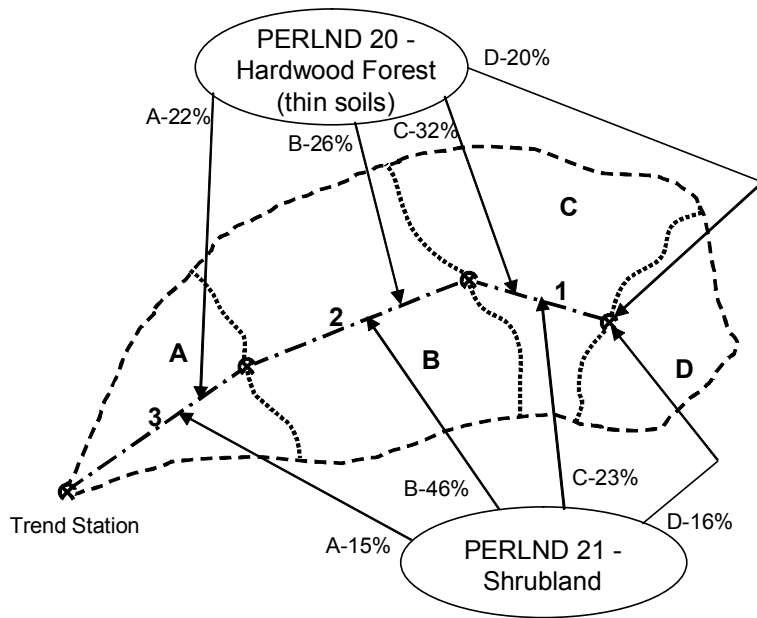


Figure 7 : A Comparison of Hydrographs for the Simulation of Twelve Pole Creek

PERLND Grouping Within the HSPF-CHIA Model

Within the HSPF-CHIA model, the grouping approach to modeling each PERLND (one for each land use/cover classification) was selected since it accumulates all areas of like land use/cover classification within the watershed into a single PERLND. This effectively reduces model complexity and the number of parameters that must be calibrated. Figure 8 illustrates the principle behind the distribution of PERLND outflows based on the percent area of its land use/cover classification contained within each sub-watershed. Each sub-watershed has a single stream segment (RCHRES) to which its outflow is assigned. Each PERLND outflow to a particular stream segment is based on the fraction of its land use/cover classification area contained in the contributing sub-watershed.

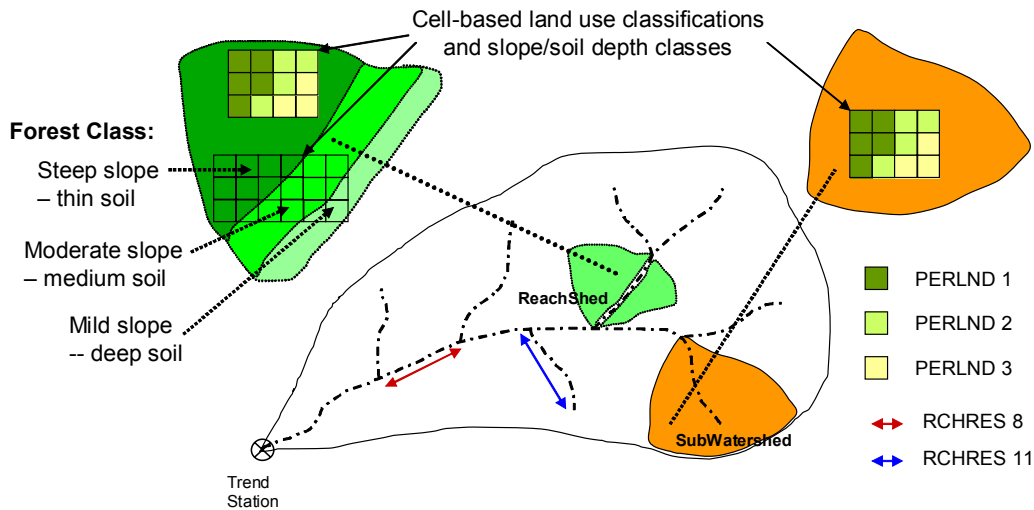


Group option for PERLND definition

Figure 8 : Grouping Land Use/Cover Classifications across Sub-watershed Boundaries.

Implementation of Land Use/Cover Classifications in PERLND Grouping

Figure 9 illustrates how the 11 different land use/cover classifications selected for the HSPF-CHIA model are implemented. Since the Forest classification is by far the most prevalent on each trend station watershed, it is subdivided into three slope categories, steep, moderate, and mild. The remaining 8 categories are not subdivided by slope, since their portion of the watershed area is typically a small percentage. Preliminary calibration experience seemed to point out a need to provide slope differentiation in the most prevailing classification, since it was logical to assume that there are significant hydrologic response differences between steep and milder slopes for the forest classification. The forest data slope categories were computed using the underlying DEM, and then incorporated into the land use/cover classification GIS layer, which is based on the 1993 GAP data (Strager and Yuill, 2002). Each grid cell is classified according to one of the 11 assigned land use/cover classifications. Within each sub-watershed the area associated with each classification is assigned to its corresponding PERLND, and a record is maintained of which stream segment receives the outflow from that area (see Figure 8).



Segmentation Based on NHD Stream Drainage Network Representation in WCMS.

Figure 9 : Illustration of Assignment of Land Use/Cover Classification in PERLND Grouping.

Manual Calibration and Verification Using HSPEXP

In order to begin the HSPF-CHIA calibration, initial values of selected calibration parameters needed to be assigned. These initial values were based on a review of parameters from other calibration studies within the Mid-Atlantic region, as determined from the HSPFParm, (1999) database (a database maintained by EPA as part of the BASINS software package), and values from similar studies (Sams, et al., (1995)), including EPA BASINS Technical Note 6, (2000). Personal communications with Kate Flynn of the USGS, Reston, in 2003 resulted in a calibration procedure that uses a single HSPF uci that is designed to combine all of the calibration watersheds into a single HSPF model run. Following a combined HSPF model run, the current calibration parameters could then be checked for suitability using a utility program called HSPEXP (USGS Report 94-4168, (1994)). This approach resulted in the creation of a single uci for Twelve Pole Creek, Buffalo Creek, Tygart Valley at Elkins, Clear Fork, and Big Sandy (Figure 1). Some simplifications were required since HSPEXP has a limit on the number of PERLND's and RCHRES's it can handle at one time, which is the reason for the 600 hectare threshold watershed subdivision used for calibration (Figure 6). Successful HSPF calibration runs were made using the combined uci within the HSPEXP software. A second combined uci was created for the 5 verification watersheds: Panther Creek, Piney Creek, Tygart Valley at Belington, Tygart Valley at Daily, and Big Sandy (Figure 2).

Table 3 shows the manual calibration results for the calibration watersheds, while Table 4 shows the corresponding results for the verification watersheds. The performance of the model is evaluated in HSPEXP by a number of statistics that are included in both tables. The statistics are based on average annual values, and show that, in most cases, the total runoff depths in each of the categories are in good agreement. The data available for calibration is considered the bare minimum for HSPF applications; therefore, it was impossible to meet the standard error criteria limits in all cases. However, since the application of HSPF for CHIA is a comparative analysis between the baseline hydrology and water quality, to that following additional mining, absolute accuracy is less important than comparative accuracy. The calibration errors are considered acceptable for the needs of CHIA, when used in the comparative analysis mode.

To provide additional calibration confidence, a detailed model performance analysis and parameter optimization study was conducted using independent optimization software (see following section).

Table 3 : HSPF Model Calibration Statistics, Simulation Period: 1/1/1985-1/1/1990.

	TWELVE POLE CREEK		BUFFALO CREEK		TYGART VALLEY AT ELKINS	
	Simulated	Observed	Simulated	Observed	Simulated	Observed
Total runoff, in inches	107.3	102.9	123.9	121.21	179.4	167.544
Total of highest 10% flows, in inches	58.25	57.3	66.7	63.11	89.22	78.442
Total of lowest 50% flows, in inches	6.91	5.39	9.58	10.33	17.98	18.503
	Simulated	Potential	Simulated	Potential	Simulated	Potential
Evapotranspiration, in inches	123.7	131.8	148.5	153.7	140.1	142.2
	Simulated	Observed	Simulated	Observed	Simulated	Observed
Baseflow recession rate	0.92	0.91	0.91	0.92	0.9	0.91
Summer flow volume, in inches	8.39	6.68	10.4	13.13	23.97	22.74
Winter flow volume, in inches	50.67	49.33	46.14	46.6	54.1	53.89
	Current	Criteria	Current	Criteria	Current	Criteria
Error in total volume	4.3	10	2.2	10	7.1	10
Error in low flow recession	-0.01	0.01	0.01	0.01	0.01	0.01
Error in 50% lowest flows	28.3	10	-7.2	10	-2.8	10
Error in 10% highest flows	1.7	15	5.7	15	13.7	15
Seasonal volume error	19.7	10	19.8	15	5	10

	CLEAR FORK		BIG SANDY	
	Simulated	Observed	Simulated	Observed
Total runoff, in inches	113.9	109.298	179	173.25
Total of highest 10% flows, in inches	58.27	54.427	86.62	76.686
Total of lowest 50% flows, in inches	8.75	9.451	19.37	21.924
	Simulated	Potential	Simulated	Potential
Evapotranspiration, in inches	149.7	154.6	118.7	120
	Simulated	Observed	Simulated	Observed
Baseflow recession rate	0.91	0.93	0.92	0.92
Summer flow volume, in inches	6.76	6.904	20.55	19.729
Winter flow volume, in inches	43.56	44.658	51.44	63.8
	Current	Criteria	Current	Criteria
Error in total volume	4.2	10	3.7	10
Error in low flow recession	0.02	0.01	0	0.01
Error in 50% lowest flows	-7.4	10	-11.6	10
Error in 10% highest flows	7.1	15	13	15
Seasonal volume error	0.4	10	23.6	10

Table 4 : HSPF Model Verification Statistics, Simulation Period: 1/1/1976-12/31/1981.

	TYGART VALLEY BELINGTON		PINEY CREEK		PANTHER CREEK	
	Simulated	Observed	Simulated	Observed	Simulated	Observed
Total runoff, in inches	149.9	162.593	114.7	90.047	102.4	102.117
Total of highest 10% flows, in inches	66.42	66.21	56.48	35.638	56.51	54.585
Total of lowest 50% flows, in inches	14.48	20.134	10.53	12.745	7.05	8.586
	Simulated	Potential	Simulated	Potential	Simulated	Potential
Evapotranspiration, in inches	113.3	114.5	104.1	108.6	134.8	137.5
	Simulated	Observed	Simulated	Observed	Simulated	Observed
Baseflow recession rate	0.87	0.91	0.88	0.92	0.9	0.9
Summer flow volume, in inches	15.18	21.151	10.01	11.037	5.36	8.489
Winter flow volume, in inches	60.87	63.773	43.04	33.556	39.84	40.357
	Current	Criteria	Current	Criteria	Current	Criteria
Error in total volume	-7.8	10	27.4	10	0.3	10
Error in low flow recession	0.04	0.01	0.04	0.01	0	0.01
Error in 50% lowest flows	-28.1	10	-17.4	10	-17.9	10
Error in 10% highest flows	0.3	15	58.5	15	3.5	15
Seasonal volume error	23.6	10	37.6	10	35.6	10

	BIG SANDY		TYGART VALLEY DAILEY	
	Simulated	Observed	Simulated	Observed
Total runoff, in inches	147.6	163.32	158.9	157.525
Total of highest 10% flows, in inches	79.67	66.886	72.69	66.621
Total of lowest 50% flows, in inches	11.12	19.971	14.4	18.782
	Simulated	Potential	Simulated	Potential
Evapotranspiration, in inches	92.51	93.3	109.6	110.3
	Simulated	Observed	Simulated	Observed
Baseflow recession rate	0.9	0.91	0.88	0.9
Summer flow volume, in inches	15.31	21.403	15.85	20.686
Winter flow volume, in inches	55.42	62.554	63.75	63.183
	Current	Criteria	Current	Criteria
Error in total volume	-9.6	10	0.9	10
Error in low flow recession	0.01	0.01	0.02	0.01
Error in 50% lowest flows	-44.3	10	-23.3	10
Error in 10% highest flows	19.1	15	9.1	15
Seasonal volume error	17.1	10	24.3	10

HSPF-CHIA Calibration Performance Evaluation and Optimization Study

Performance Evaluation Procedures

There are a number of publications dealing with the evaluation of watershed model performance. Although no uniform criteria have been established, the general view is that it is advisable to report several criteria in order to more objectively quantify the performance of a given hydrological model. Following criteria were used to determine an optimum set of HSPF parameters and to estimate the predictive ability of the model.

The first two, the Deviation of Runoff Volumes, D_V , and the Coefficient of Efficiency (known also as the Nash-Sutcliffe coefficient), E , are suggested by the ASCE Task Committee on Definition of Criteria for Evaluation of Watershed Models (ASCE, 1993):

Deviation of Runoff Volume, D_V (Martinec and Rango, 1989)

$$D_V = \frac{V - V_S}{V} \quad (1)$$

where: V – the total observed runoff volume for the simulation period
 V_S – the total simulated runoff volume for the simulation period.

It should be noted that for a perfect model, D_v equals zero. A hydrological model is considered very good if $D_v < 0.1$, good if $0.1 < D_v < 0.15$ and fair if $0.15 < D_v < 0.25$.

Coefficient of Efficiency, E (Nash and Sutcliffe, 1970)

$$E = 1.0 - \frac{\sum_{i=1}^N (O_i - S_i)^2}{\sum_{i=1}^N (O_i - \bar{O})^2} \quad (2)$$

where: O_i – the observed daily discharge
 S_i – the model simulated daily discharge
 \bar{O} – the average observed discharge
 N – the number of discharges values.

The Coefficient of Efficiency is the ratio of the Mean Square Error (MSE) to the variance in the observed data subtracted from unity and it ranges from unity (ideal model) to minus infinity (poor model). When $E = 0$ the square differences between the model simulation and the observation is equal to the variability in the observed data, which means that the observed mean, \bar{O} , is as good a predictor as the model.

Legates and McCabe (1999) suggested using several other criteria that would give a better representation of the efficacy of model simulations and recommended a wide range of statistics to be reported including the observed and simulated means and standard deviations.

Modified Coefficient of Efficiency, $E1$ (Legates and McCabe, 1999)

$$E1 = 1.0 - \frac{\sum_{i=1}^N |O_i - S_i|}{\sum_{i=1}^N |O_i - \bar{O}|} \quad (3)$$

The Modified Coefficient of Efficiency has a meaning similar to the Coefficient of Efficiency but it is considered preferable because it reduces the effect of squaring in statistics by giving errors and differences their appropriate weight.

In a similar way, the Index of Agreement (Willmott, 1981), another descriptive relative error measure that reflects agreement between simulated and observed values, was proposed in adjusted form, namely the Modified Index of Agreement, $d1$, in order to eliminate squared values of errors.

Modified Index of Agreement, $d1$ (Legates and McCabe, 1999)

$$d1 = 1.0 - \frac{\sum_{i=1}^N |O_i - S_i|}{\sum_{i=1}^N (|S_i - \bar{O}| + |O_i - \bar{O}|)} \quad (4)$$

The Modified Index of Agreement varies from 0.0 for a full disagreement of simulated and observed values to 1.0 for a perfect model.

The use of two non-negative statistics is also recommended in order to arrive at a more comprehensive model evaluation (Willmott, 1984 and Legates, 1999). These absolute error measures (the error in the units of variable) are the Mean Absolute Error (*MAE*) and the Square-root of the Mean Square Error (*RMSE*).

Mean Absolute Error, *MAE* (Willmott, 1984)

$$MAE = N^{-1} \left[\sum_{i=1}^N |O_i - S_i| \right] \quad (5)$$

Compared to *RMSE*, the *MAE* is less sensitive to extreme values. Generally, *RMSE* is bigger than *MAE* and, thus, the relative degree of difference between them represents the extent of the variance of absolute errors.

Square-root of the Mean Square error, *RMSE* (Willmott, 1984)

$$RMSE = \left[N^{-1} \sum_{i=1}^N (O_i - S_i)^2 \right]^{0.5} \quad (6)$$

The advantage of *RMSE* is that it provides valuable information about sources of error. Willmott (1981) recommended portioning *RMSE* into “systematic” (*RMSE_s*) and “unsystematic” (*RMSE_u*) error components in order to determine the nature of error. The relationship is expressed in the following equation (Willmott, 1984):

$$RMSE^2 = RMSE_s^2 + RMSE_u^2 \quad (7)$$

RMSE_s is actually the bias or deviation of the linear regression line slope on an observed versus simulated plot from a 45 deg. line and, therefore, it indicates the ability of a model to replicate variations in observed data. *RMSE_u* is a component responsible for random variations.

“Systematic” portion of the *RMSE*, *RMSE_s* (Willmott, 1984)

$$RMSE_s = \left[N^{-1} \sum_{i=1}^N (\bar{S}_i - O_i)^2 \right]^{0.5} \quad (8)$$

and

“Unsystematic” portion of the *RMSE*, *RMSE_u* (Willmott, 1984)

$$RMSE_u = \left[N^{-1} \sum_{i=1}^N (S_i - \bar{S}_i)^2 \right]^{0.5} \quad (9)$$

where: $\bar{S}_i = a + bO_i$

a and b are parameters of a linear regression between observed and simulated values. For a good model performance, $RMSE_s$ should be low and $RMSE_u$ should approach $RMSE$, which itself should be low.

Another widely used statistic is the Coefficient of Determination, R^2 . Although it has been argued by Willmott (1984) and Legates (1999) that the Coefficient of Determination has a lack of sensitivity to additive and proportional differences between observed and model simulated values, and is sensitive to outliers (extreme values), it remains a quite popular and commonly reported criterion.

Coefficient of Determination, R^2 (Legates and McCabe, 1999)

$$R^2 = \frac{\left[\sum_{i=1}^N (O_i - \bar{O})(S_i - \bar{S}) \right]}{\left[\sum_{i=1}^N (O_i - \bar{O})^2 \right]^{0.5} \left[\sum_{i=1}^N (S_i - \bar{S})^2 \right]^{0.5}} \quad (10)$$

The Coefficient of Determination ranges from 0.0 (poor model) to 1.0 (perfect model). It expresses the proportion of the total variance in the observed data that can be explained by the model.

Described criteria in combination with provided standard deviations and means and data-display graphics should be sufficient to evaluate the performance of a hydrological model (Donigian, 2002).

Optimization of the HSPF Parameters using PEST Software

As explained above, starting values of HSPF parameters for the manual calibration procedure were determined based on other studies, calibration watershed characteristics, and suggested parameter ranges in BASINS Technical Note #6. Each of eleven (11) PERLNDs was assigned an individual set of parameters. The model was then manually calibrated using the HSPEXP program. The resulting parameter sets were assumed to be the calibrated values and served as the starting point for the optimization study.

The final adjustment of parameters and calibration was preformed using the PEST optimization software program (Doherty, 2002). The PEST program adjusts specified HSPF parameters until the differences between gaged flows and the model's simulated flows are minimized according to specified multiobjective criteria.

A sensitivity study was first done to see if separation of the FOREST land-use category into three different PERLNDs according to surface slopes results in significant improvement in the performance of the model. HSPF parameters that could be influenced by surface slopes, as they are described in the BASINS Technical Note #6, were LZSN, INFILT, UZSN, INTFW, AGWRC, and IRC (note: parameter definitions are given in Bicknell, et al, 2001). PEST optimization was done twice for each of the specified parameters. In the first run one of the above parameters was set to be the same value for

all three forest PERLNDs. A second run specified that a parameter for each Forest PERLND would have independent value and not necessarily the same as those of other two. The statistical data, which are summarized in Table 5, show that slope differentiation yields no significant improvement in the performance of the model. That raises the question of the usefulness of categorizing the Forest land-use/cover by surface slope. If this differentiation were ignored, then there would be essentially a single land-use/cover classification on many of the trend station watersheds, leaving little flexibility in applying the model. Since leaving the differentiation in the model does not effect the calibration one way or another, it was decided to retain the three slope categories to allow potential future flexibility in application of the HSPF-CHIA model.

Table 5 : Selected Error Statistics Comparing a Single Forest Slope Land-use/cover Category with Subdivision Based on Slope.

TWELVEPOLE CREEK	E	E1	D _v	d1	R ²	RMSE,cfs	MAE,cfs
Initial parameters	0.42	0.38	0.17	0.70	0.71	86	33
LZSN	0.42	0.38	0.19	0.70	0.71	87	33
LZSN (three FOREST categories)	0.42	0.38	0.18	0.70	0.71	86	32
AGWRC	0.44	0.40	0.16	0.70	0.71	85	32
AGWRC (three FOREST categories)	0.44	0.40	0.16	0.70	0.71	85	32
IRC	0.45	0.38	0.16	0.69	0.70	89	33
IRC (three FORES categories)	0.45	0.38	0.16	0.69	0.70	84	33
INTFW	0.45	0.39	0.17	0.70	0.71	84	32
INTFW (three FOREST categories)	0.45	0.39	0.17	0.70	0.71	84	32
UZNS seasonal	0.51	0.42	0.14	0.71	0.73	79	31
UZNS seasonal (three FOREST cat.)	0.51	0.42	0.14	0.71	0.73	79	31
INFILT	0.52	0.41	0.15	0.69	0.73	78	31
INFILT (three FOREST categories)	0.52	0.42	0.15	0.69	0.73	78	31
BUFFALO CREEK	E	E1	D _v	d1	R ²	RMSE,cfs	MAE,cfs
Initial parameters	0.49	0.47	0.02	0.74	0.75	233	92
LZSN	0.49	0.47	0.04	0.74	0.75	238	93
LZSN (three FOREST categories)	0.49	0.47	0.02	0.74	0.75	232	92
AGWRC	0.44	0.40	0.16	0.70	0.71	235	92
AGWRC (three FOREST categories)	0.44	0.40	0.16	0.70	0.71	233	92
IRC	0.51	0.46	0.02	0.72	0.74	228	93
IRC (three FORES categories)	0.51	0.46	0.02	0.72	0.74	228	93
INTFW	0.53	0.47	0.02	0.74	0.76	224	92
INTFW (three FOREST categories)	0.53	0.47	0.02	0.74	0.76	224	92
UZSN seasonal	0.53	0.47	0.00	0.73	0.75	225	92
UZSN seasonal (three FOREST cat.)	0.53	0.47	0.00	0.73	0.75	225	92
INFILT	0.58	0.47	0.01	0.72	0.76	211	91
INFILT (three FOREST categories)	0.58	0.47	0.01	0.72	0.76	211	91
TYGART VALLEY (EIK)	E	E1	D _v	d1	R ²	RMSE,cfs	MAE,cfs
Initial parameters	0.19	0.34	0.09	0.68	0.64	819	334
LZSN	0.19	0.34	0.10	0.68	0.64	830	336
LZSN (three FOREST categories)	0.20	0.34	0.09	0.68	0.65	815	332
AGWRC	0.21	0.35	0.09	0.68	0.65	810	325
AGWRC (three FOREST categories)	0.21	0.35	0.09	0.68	0.65	810	326
IRC	0.23	0.35	0.09	0.67	0.63	799	328

IRC (three FOREST categories)	0.23	0.35	0.09	0.67	0.63	799	328
UZSN seasonal	0.26	0.37	0.08	0.69	0.66	784	315
UZSN seasonal (three FOREST cat.)	0.26	0.37	0.08	0.69	0.66	784	315
INTFW	0.31	0.38	0.09	0.69	0.68	754	314
INTFW (three FOREST categories)	0.31	0.38	0.09	0.69	0.68	761	316
INFLT	0.57	0.45	0.08	0.71	0.76	599	278
INFILT (three FOREST categories)	0.56	0.45	0.08	0.71	0.76	602	278
CLEAR FORK							
Initial parameters	E	E1	D _v	d1	R ²	RMSE,cfs	MAE,cfs
LZSN	0.46	0.42	0.07	0.72	0.74	215	97
LZSN (three FOREST categories)	0.46	0.42	0.07	0.72	0.74	216	97
LZSN (three FOREST categories)	0.47	0.42	0.07	0.72	0.74	214	96
AGWRC	0.47	0.44	0.07	0.72	0.74	213	92
AGWRC (three FOREST categories)	0.47	0.44	0.07	0.72	0.74	213	92
INTFW	0.49	0.43	0.07	0.72	0.75	209	94
INTFW (three FOREST categories)	0.49	0.43	0.07	0.72	0.75	206	96
IRC	0.51	0.43	0.07	0.71	0.74	205	93
IRC (three FORES categories)	0.51	0.43	0.07	0.71	0.74	205	93
UZSN seasonal	0.49	0.46	0.05	0.73	0.74	208	90
UZSN seasonal (three FOREST cat.)	0.49	0.46	0.05	0.73	0.74	208	90
INFILT	0.59	0.47	0.06	0.72	0.77	189	88
INFILT (three FOREST categories)	0.59	0.47	0.06	0.72	0.77	189	88
BIG SANDY							
Initial parameters	E	E1	D _v	d1	R ²	RMSE,cfs	MAE,cfs
LZSN	0.38	0.37	0.00	0.69	0.71	505	230
LZSN	0.38	0.37	0.00	0.69	0.71	512	233
LZSN (three FOREST categories)	0.38	0.37	0.00	0.69	0.71	505	230
AGWRC	0.40	0.38	0.01	0.70	0.71	500	225
AGWRC (three FOREST categories)	0.40	0.38	0.00	0.70	0.71	500	225
IRC	0.43	0.41	0.00	0.70	0.70	487	214
IRC (three FOREST categories)	0.43	0.41	0.00	0.70	0.70	491	218
UZSN seasonal	0.38	0.39	0.01	0.70	0.70	507	220
UZSN seasonal (three FOREST categories)	0.38	0.39	0.01	0.70	0.70	507	220
INTFW	0.48	0.41	0.00	0.71	0.74	462	213
INTFW (three FOREST categories)	0.48	0.41	0.00	0.71	0.74	468	216
INFILT	0.57	0.48	0.01	0.73	0.76	422	189
INFILT (three FOREST categories)	0.58	0.48	0.01	0.72	0.76	417	188

In the following calibration and sensitivity study all three PERLNDs of the Forest land-use/cover classification were relegated to the same set of parameters, thus reducing the number of calibrated sets of parameters from eleven (11) to nine (9).

A second sensitivity study was performed on each of the parameters in the set of the Forest PERLND segments. During every PEST optimization and simulation run one of these parameters was calibrated independently (the rest of parameters in a set were fixed at initially calibrated values) and statistics of the model performance were calculated. Analysis of the statistics allowed the determination of a group of parameters that were the most influential with respect to the performance of the model. The improvement in value of the Coefficient of Determination, E , and the Modified Coefficient of Determination, $E1$, for optimized parameters INFILT, UZSN, INTFW, and

IRC was more dramatic than for the other parameters. Table 6 shows the statistics for each parameter, as it alone is optimized.

Table 6 : Summary of Individual Parameter Optimization Statistics

TWELVEPOLE CREEK	E	E1	D _v	d1	R ²	RMSE (cfs)	MAE (cfs)
Initial parameters	0.42	0.38	0.17	0.70	0.71	85.91	32.63
LZSN	0.41	0.37	0.19	0.70	0.71	87.10	33.16
AGWETP	0.42	0.38	0.16	0.70	0.71	85.96	32.64
LZETP	0.42	0.38	0.17	0.70	0.71	85.91	32.63
DEEPFR	0.43	0.39	0.10	0.70	0.71	85.44	31.97
BASETP	0.43	0.38	0.17	0.70	0.71	85.62	32.48
AGWRC	0.44	0.40	0.16	0.70	0.71	84.87	31.65
IRC	0.45	0.38	0.16	0.69	0.70	83.99	32.67
INTFW	0.45	0.39	0.17	0.70	0.71	83.80	32.03
UZNS seasonal	0.51	0.42	0.14	0.71	0.73	78.90	30.72
INFILT	0.52	0.41	0.15	0.69	0.73	78.14	30.74
BUFFALO CREEK	E	E1	D _v	d1	R ²	RMSE (cfs)	MAE (cfs)
Initial parameters	0.49	0.47	0.02	0.74	0.75	233.41	92.12
LZSN	0.47	0.46	0.04	0.74	0.75	237.51	92.96
AGWETP	0.49	0.47	0.02	0.74	0.75	233.61	92.22
LZETP	0.49	0.47	0.02	0.74	0.75	233.41	92.12
DEEPFR	0.49	0.46	0.05	0.74	0.75	232.77	92.50
BASETP	0.49	0.47	0.02	0.74	0.75	232.31	91.55
AGWRC	0.50	0.48	0.01	0.74	0.75	230.61	89.35
IRC	0.51	0.46	0.02	0.72	0.74	227.84	92.87
INTFW	0.53	0.47	0.02	0.74	0.76	223.82	91.54
UZSN seasonal	0.53	0.47	0.00	0.73	0.75	224.51	91.81
INFILT	0.58	0.47	0.01	0.72	0.76	211.15	91.02
TYGART VALLEY (EIk)	E	E1	D _v	d1	R ²	RMSE (cfs)	MAE (cfs)
Initial parameters	0.19	0.34	0.09	0.68	0.64	819.49	333.74
LZSN	0.17	0.33	0.10	0.68	0.64	829.87	335.68
AGWETP	0.19	0.34	0.09	0.68	0.64	819.90	333.86
LZETP	0.19	0.34	0.09	0.68	0.64	819.49	333.74
DEEPFR	0.20	0.34	0.03	0.68	0.64	817.63	334.43
BASETP	0.20	0.34	0.09	0.68	0.65	814.80	331.79
AGWRC	0.21	0.35	0.09	0.68	0.65	809.71	325.25
IRC	0.23	0.35	0.09	0.67	0.63	798.99	328.13
UZSN seasonal	0.26	0.37	0.08	0.69	0.66	784.24	315.50
INTFW	0.31	0.38	0.09	0.69	0.68	754.30	313.57
INFLT	0.57	0.45	0.08	0.71	0.76	598.78	278.10
CLEAR FORK	E	E1	D _v	d1	R ²	RMSE (cfs)	MAE (cfs)
Initial parameters	0.46	0.42	0.07	0.72	0.74	215.47	96.65
LZSN	0.45	0.41	0.07	0.72	0.74	225.16	96.94
AGWETP	0.46	0.41	0.07	0.72	0.74	215.65	96.84
LZETP	0.46	0.42	0.07	0.72	0.74	215.47	96.65
DEEPFR	0.46	0.42	0.01	0.72	0.74	214.58	96.54
BASETP	0.47	0.42	0.07	0.72	0.74	214.22	95.75

AGWRC	0.47	0.44	0.07	0.72	0.74	212.82	92.47
INTFW	0.49	0.43	0.07	0.72	0.75	208.71	94.27
UZSN seasonal	0.49	0.46	0.05	0.73	0.74	208.27	89.96
IRC	0.51	0.43	0.07	0.71	0.74	205.35	93.43
INFILT	0.59	0.47	0.06	0.72	0.77	188.56	87.52
BIG SANDY Initial parameters	E	E1	D _v	d1	R ²	RMSE (cfs)	MAE (cfs)
	0.38	0.37	0.00	0.69	0.71	505.24	229.67
LZSN	0.37	0.36	0.01	0.69	0.71	512.18	232.80
AGWETP	0.38	0.37	0.00	0.69	0.71	505.50	229.84
LZETP	0.38	0.37	0.00	0.69	0.71	505.24	229.67
DEEPFR	0.38	0.35	0.06	0.69	0.70	506.13	234.78
BASETP	0.38	0.37	0.00	0.69	0.71	503.27	228.01
UZSN seasonal	0.38	0.39	0.01	0.70	0.70	506.67	219.97
AGWRC	0.40	0.38	0.00	0.70	0.71	499.81	225.36
IRC	0.43	0.41	0.00	0.70	0.70	486.69	213.69
INTFW	0.48	0.41	0.00	0.71	0.74	461.64	212.58
INFILTR	0.57	0.48	0.01	0.73	0.76	422.38	189.14

In a next step of the calibration and sensitivity analysis, the PEST program was setup to optimize all eleven parameters for each of nine (9) PERLNDs in the same run (the Forest Land categories were constrained to use an identical parameter set). The final optimized parameter values are presented in Table 7 and the statistical data from this run are shown in Table 9 as run # 1. Although the deviations of runoff volume, D_v , predominately show a good model performance, values of Coefficient of Efficiency, E , are not very high for the calibration watersheds and somewhat lower for the verification watersheds.

Table 7 : PEST Optimized Parameters Values for 11 PERLND Segments.

#	PERLND SEGMENT	PARAMETERS										
		LZSN (in)	INFILT (in/hr)	KVARY (1/in)	AGWCR (1/day)	DEEPFR	BASETP	AGWETP	UZSN (in)	INTFW	IRC (1/day)	LZETP
1	Forest Land Steep	4.99	0.012	0.001	0.938	0.21	0.02	0.0099	0.499	3.5	0.27	0.599
2	Forest Land Moderate	4.99	0.012	0.001	0.938	0.21	0.02	0.0099	0.499	3.5	0.27	0.599
3	Forest Land Mild	4.99	0.012	0.001	0.938	0.21	0.02	0.0099	0.499	3.5	0.27	0.599
4	Pasture/ Grassland	4.99	0.020	0.0010	0.938	0.21	0.02	0.0099	0.704	3.5	0.27	0.462
5	Urban/ Developed	4.86	0.004	0.0090	0.933	0.21	0.02	0.0099	0.868	4.2	0.28	0.205
6	Mined Land	6.93	0.020	0.0020	0.934	0.21	0.02	0.0099	0.502	3.5	0.26	0.209
7	Barren Land	5.71	0.014	0.0190	0.971	0.21	0.02	0.0099	0.271	5.0	0.30	0.453
8	Surface Water	4.33	0.037	0.0010	0.912	0.21	0.02	0.0099	0.974	2.1	0.24	0.115
9	Row Crop Agricuilt.	4.97	0.020	0.0060	0.938	0.21	0.02	0.0099	0.500	3.7	0.26	0.319
10	Wetland	6.99	0.007	0.0060	0.953	0.21	0.02	0.0099	0.967	3.8	0.26	0.501

The final PEST calibration was done using the parameters INFILT, UZSN, INTFW, and IRC as variables, while holding the remaining parameters fixed at their original values within the Forest land-use/cover PERLND. These parameters are those identified earlier to be the most influential in the sensitivity study. Forest is the major land-use/cover for all the calibration watersheds, and therefore, its parameters will be the most influential. Table 8 list the values of the optimized parameters and compares them to values obtained by other investigators in unrelated calibration studies.

The analyzed statistics, which are summarized in Table 9 as a run # 2, show a good model performance with the Coefficient of Efficiency, *E*, the possible range of which is from minus unity (poor model) to 1.0 (perfect model), at 0.57 for the entire simulation period of six years for 5 calibration watersheds. This is in a range of values reported by similar investigations. The deviation of total runoff volume for four of the watersheds is less then 15% which is rated as good performance.

Table 8 : Comparison of the Final Set of Parameters for the Forest PERLND categories with those reported by other studies.

Parameter	Units	Optimized value	Moore et al. 1988	Chew et al. 1991	Laroche et al. 1996	Srinivasan et al. 1999	Engelmann 1999	Doherty et al. 2003
LZSN	in	3	4.9	5	14	5.12-5.9	3.82	2-2.58
INFILT	in/hr	0.0634	0.004-0.0196	0.063-0.14	0.23	0.039-0.39	0.0394	0.028-0.071
UZSN	in	0.96 ^S	0.197	0.016-0.043	0.756	0.31-0.75	0.7	1.55-2
KVARY	in ⁻¹	3	----	----	0.06	----	0.61	----
AGWRC	day ⁻¹	0.983	0.98	----	0.99	0.8-0.9	0.99	0.942-0.988
INFEXP		2	-----	----	0	----	2	----
INTFW		2.62	1	----	9.83	0.8-3.0	0.5	1-1.31
INFILD		2*	----	----	1.99	----	1	-----
IRC	day ⁻¹	0.5	0.1	-----	0	----	-----	0.499-0.533
LZETP		0.5 ^S	0.3-0.55	----	0-0.8	0.1-0.8	0.42	0.5
DEEPPFR		0.2	----	----	----	----	0.1	0.1(fixed)
BASETP		0.05	----	----	----	----	0.02	0.157-0.182
AGWETP		0.01*	-----	----	-----	----	0	0.02-0.027

* Fixed initial value

--- Not mentioned in reference

^S Monthly variable value (table below)

	Jan	Feb	Mar	April	May	June	July	Aug	Sept	Oct	Nov	Dec
UZSN	1.66	1.54	1.28	0.92	0.56	0.33	0.26	0.38	0.66	1.01	1.36	1.59
LZETP	0.056	0.006	0.089	0.28	0.53	0.78	0.95	0.99	0.9	0.73	0.47	0.23

The verification study used the five watersheds shown in Figure 2. The precipitation data were taken from the five year period between 1971 and 1975. The HSPF runs were performed using two sets of parameters that were the results of optimization runs #1 and #2, described above. A summary of statistics for the verification watersheds is included in Table 9.

Conclusions of the HSPF Model Performance Evaluation

Two sets of HSPF parameters were developed by means of PEST optimization runs. An additional set was obtained by a manual calibration by using the HSPEXP. These sets are defined and presented in Table 9:

Run # 1 – All parameters for nine (9) PERLND segments were optimized jointly using PEST.

Run # 2 – Optimization was done for the most sensitive parameters and only for Forest land-use/cover with all slope categories treated as equivalent.

Run # 3 – All parameters for eleven (11) PERLNDs were manually calibrated with the guidance of the HSPFEXP program.

A comparison of the statistics of the HSPF model performances for three runs above includes the means and standard deviations of the predicted and actual discharges during the period of the calibration. Run #1 has an average error of less than 10 % for all calibration watersheds and three of the verification watersheds. For Run #2, the simulated mean is less than 12% for calibration watersheds and less than 15 % for four of the verification watersheds. Statistics for Run # 3 indicate less than 13 % average error for the calibration watersheds and less than 13% for four of the verification watersheds. For all three runs, verification watersheds Piney and Big Sandy (1971-1975) show the highest average error, 13-22%. Comparison of observed and predicted standard deviations demonstrates that all three runs do not give a good description of observed flow variance, with an especially poor performance noted on Tygart (Elkins) and Big Sandy calibration watersheds, and Tygart (Belington), Tygart (Dailey) and Big Sandy verification watersheds.

Deviations of total runoff volume (D_v) over the entire simulation period for every run gives an uneven picture, with the model performance varying from very good to just fair, depending on the watershed. For example, for the Big Sandy watershed all three runs indicate a good model performance with D_v less than 10 percent for the simulation period, but for the verification period for the same watershed, the performance can be classified only as fair. Generally, averaging performances of all watersheds involved (calibration and verification), run # 3 produced the smallest runoff deviations and run # 2 the largest.

For years, correlation-based statistics such as the Coefficient of Determination, R^2 , which describes the degree of collinearity between model predictions and observations, were the most popular measurements of demonstrating the accuracy of a hydrologic model. Although it was pointed out by some investigations (Willmott, 1981; Willmott, 1984; Willmott et al., 1985; Legates and McCabe, 1999) that these kinds of measurements have limitations and may mislead a model evaluation, these statistics remain widely used and reported. Donigian (2002), based on his many years of experience working with the HSPF, derived a chart (see Figure 10) that evaluates model performance by ranges of values of the Coefficient of Determination. According to this table, run # 2 outperformed the other two runs. By Donigian's standards, run # 2 demonstrated good performance for calibration watersheds and very good performance for verification watersheds.

Slight differences in the Modified Index of Agreement, in run # 2, that produced slightly higher values as compared to the other runs are of minimal use in determining best performance. However, both the Coefficient of Efficiency, E , and the Modified

Coefficient of Efficiency, $E1$, which are suggested as the most appropriate relative error available (Legates, 1999) show better (higher) values for run # 2. Also, E and $E1$ of run # 2 are more consistent for all considered watersheds. Runs # 1 and #3 produce negative E for the verification Piney Creek watershed, which indicates that the observed mean is a better predictor than the model.

At first glance the magnitudes of absolute error statistics MAE and $RMSE$ are lower for run # 2. However, further segmentation of $RMSE$ into systematic and unsystematic sources of error presents a different picture with runs #1 and #3 having a smaller systematic and higher unsystematic error values than run # 2, and therefore, they have a higher potential accuracy.

The conclusion, with regard to selection of the best parameter set, of the three sets compared above, remains to be determined at the time of this report. Overall, runs #2 and #3 have positive features, with run #2 yielding the best statistical results. Currently, the run #2 parameter set is being used in the WCMS-HSPF model.

Table 9. Summary of Calibration Run Statistics for Calibration and Verification Watersheds

Calibration	wsheds	1	2	3	4	5	6	7	8	9	10	11	12	13
Mean Observed \bar{O} , cfs		Mean Simulated \bar{S} , cfs	Standard Deviation of O , cfs S_o	Standard Deviation of S , cfs S_s	Coef. of Efficiency E	Coef. of Efficiency E_1	Modified Coef. of Efficiency E_1	Deviation of Runoff Volume D_v	Modif. index of agr. d1	Coef. of Determination R^2	Root m. sq. err. cfs RMSE	Systematic RMSE, cfs RMSEs	Unsystematic RMSE, cfs RMSEu	Mean abs --- cfs MAE
Twelvepole	run #1	49.23	113.24	120.69	0.34	0.33	0.33	0.11	0.69	0.69	92.03	30.05	86.96	35.20
Twelvepole	run #2	49.23	113.24	72.85	0.57	0.48	0.48	0.01	0.72	0.76	74.56	57.72	47.17	27.30
Twelvepole	run #3	49.23	113.24	110.82	0.40	0.37	0.37	0.13	0.69	0.69	87.86	36.84	79.75	32.94
Buffalo	run #1	173.61	326.12	363.89	0.42	0.42	0.42	0.04	0.72	0.75	248.24	54.80	242.06	100.82
Buffalo	run #2	152.90	326.12	228.81	0.55	0.55	0.55	0.12	0.72	0.75	218.26	156.51	152.05	90.50
Buffalo	run #3	173.61	326.12	349.45	0.39	0.44	0.44	0.00	0.72	0.73	249.98	67.67	240.61	95.83
Tygart(E)	run #1	553.09	911.35	755.43	0.61	0.50	0.50	0.02	0.74	0.78	566.86	318.88	468.49	249.91
Tygart(E)	run #2	553.09	911.35	756.02	0.60	0.50	0.50	0.03	0.74	0.78	576.45	324.54	476.23	250.69
Tygart(E)	run #3	553.09	911.35	1078.79	0.00	0.27	0.27	0.08	0.65	0.60	905.76	266.76	865.37	364.00
Clear Fork	run #1	171.68	292.94	333.91	0.30	0.32	0.32	0.01	0.68	0.70	245.70	59.22	238.40	112.64
Clear Fork	run #2	153.29	292.94	223.53	0.58	0.54	0.54	0.10	0.75	0.76	189.48	123.07	144.02	76.36
Clear Fork	run #3	174.55	292.29	298.47	0.43	0.39	0.39	0.03	0.71	0.72	220.75	73.70	206.78	99.87
Big Sandy	run #1	427.52	643.19	727.43	0.29	0.28	0.28	0.04	0.67	0.69	541.94	139.51	523.55	261.71
Big Sandy	run #2	381.16	643.19	509.35	0.55	0.49	0.49	0.11	0.74	0.75	432.39	267.56	339.53	184.24
Big Sandy	run #3	427.52	643.19	705.09	0.28	0.32	0.32	0.00	0.68	0.68	541.53	163.73	517.82	243.57
Verification	wsheds	1	2	3	4	5	6	7	8	9	10	11	12	13
Tygart(B)	run #1	1041.93	1380.20	1486.69	0.35	0.31	0.31	0.10	0.67	0.70	1115.79	352.03	1058.45	596.28
Tygart(B)	run #2	1041.93	1380.20	1071.42	0.67	0.54	0.54	0.14	0.75	0.83	790.26	513.71	600.17	399.84
Tygart(B)	run #3	1041.93	1380.20	1459.52	0.28	0.32	0.32	0.07	0.67	0.66	1173.43	421.31	1094.79	584.71
Piney	run #1	74.86	100.44	153.74	-0.11	0.11	0.11	0.16	0.63	0.73	106.03	52.22	104.60	17.10
Piney	run #2	74.86	100.44	112.73	0.55	0.38	0.38	0.13	0.71	0.81	67.64	13.56	66.25	36.24
Piney	run #3	74.86	100.44	154.27	-0.22	0.14	0.14	0.22	0.64	0.71	110.80	18.49	109.21	50.54
Panther	run #1	49.95	97.27	108.30	0.61	0.46	0.46	0.06	0.75	0.83	60.54	7.81	60.02	27.03
Panther	run #2	49.95	97.27	71.01	0.67	0.56	0.56	0.13	0.77	0.83	55.82	39.04	39.87	21.92
Panther	run #3	49.95	97.27	96.51	0.68	0.53	0.53	0.05	0.77	0.84	54.81	16.33	52.30	23.41
Big Sandy	run #1	515.41	726.86	907.71	0.14	0.24	0.24	0.15	0.66	0.68	675.93	132.14	662.69	323.16
Big Sandy	run #2	515.41	726.86	667.51	0.45	0.40	0.40	0.20	0.71	0.71	539.30	269.52	466.96	252.77
Big Sandy	run #3	515.41	726.86	890.30	0.09	0.25	0.25	0.13	0.66	0.65	694.92	161.40	675.71	315.96
Tygart(D)	run #1	459.87	660.61	726.87	0.40	0.32	0.32	0.02	0.68	0.73	513.59	130.32	496.63	266.34
Tygart(D)	run #2	459.87	660.61	508.65	0.70	0.55	0.55	0.07	0.76	0.84	363.68	235.81	276.75	176.55
Tygart(D)	run #3	459.87	660.61	714.87	0.33	0.33	0.33	0.01	0.68	0.70	538.80	163.23	513.31	260.96

Figure 10 : R and R2 Value Ranges for Model Performance (Donigian, A., 2002)

Criteria				
R	← 0.75	0.80	0.85	0.90 0.95 →
R ²	← 0.6	0.7	0.8	0.9 →
Daily Flows	Poor	Fair	Good	Very Good
Monthly Flows	Poor	Fair	Good	Very Good

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Appendix B. Final Calibration User's Control Input File (UCI)

RUN

GLOBAL

```
UCI Created by WinHSPF for Combined
START      1985/01/02 01:00  END      1990/12/31 23:00
RUN INTERP OUTPT LEVELS    1    2
RESUME     0 RUN           1
                                UNITS      1
END GLOBAL
```

FILES

```
<FILE> <UN#>***<----FILE NAME----->
>
MESSU      24  Combined.ech
           91  Combined.out
WDM        25  Combined.wdm
END FILES
```

OPN SEQUENCE

```
INGRP                                INDELT 01:00
```

```
*** TWELVE POLE
```

```
PERLND    101
PERLND    111
PERLND    121
PERLND    201
PERLND    301
PERLND    401
PERLND    601
PERLND    701
PERLND    801
RCHRES     3
RCHRES     2
RCHRES     1
```

```
*** BUFFALO
```

```
PERLND    102
PERLND    112
PERLND    122
PERLND    202
PERLND    302
PERLND    402
PERLND    602
PERLND    702
PERLND    802
PERLND    902
RCHRES     5
RCHRES     6
RCHRES     4
RCHRES     7
RCHRES     8
```

```
*** TYGART AT ELKINS
```

```
PERLND    103
PERLND    113
PERLND    123
PERLND    203
PERLND    303
PERLND    403
PERLND    603
PERLND    703
PERLND    803
PERLND    903
RCHRES    11
```



```

RCHRES      13
RCHRES      14
RCHRES      15
RCHRES      12
RCHRES      10
RCHRES       9

```

```

***          CLEAR FORK

```

```

PERLND      104
PERLND      114
PERLND      124
PERLND      204
PERLND      304
PERLND      404
PERLND      604
PERLND      704
PERLND      804
PERLND      904
RCHRES      16
RCHRES      17
RCHRES      18
RCHRES      19
RCHRES      20

```

```

***          BIG SANDY

```

```

PERLND      205
PERLND      805
PERLND      305
PERLND      905
PERLND      705
PERLND      405
PERLND      125
PERLND      115
PERLND      105
PERLND      605
RCHRES      22
RCHRES      21
RCHRES      23
RCHRES      25
RCHRES      24
RCHRES      26
RCHRES      27
COPY        1
COPY        2
COPY        3
COPY        4
COPY        5

```

```

END INGRP
END OPN SEQUENCE

```

```

PERLND

```

```

ACTIVITY
*** <PLS >          Active Sections          ***
*** x - x ATMP SNOW PWAT SED PST PWG PQAL MSTL PEST NITR PHOS TRAC ***
101 905 1 1 1 0 0 0 0 0 0 0 0 0
END ACTIVITY

```

```

PRINT-INFO
*** < PLS>          Print-flags          PIVL
PYR
*** x - x ATMP SNOW PWAT SED PST PWG PQAL MSTL PEST NITR PHOS TRAC
101 6 4 6 6 6 6 6 6 6 6 6 6 6 1
12
102 905 6 6 6 6 6 6 6 6 6 6 6 6 1
12

```

END PRINT-INFO

GEN-INFO

```
***
*** <PLS >
*** x - x
***
          TWELVE POLE CREEK
101 105Forest Land Steep      1 1 91 0 0 0
111 115Forest Land Moderate  1 1 91 0 0 0
121 125Forest Land Mild     1 1 91 0 0 0
201 205Pasture/Grassland     1 1 91 0 0 0
301 305Urban/Developed      1 1 91 0 0 0
401 405Mined Land           1 1 91 0 0 0
601 605Barren Land          1 1 91 0 0 0
701 705Surface Water        1 1 91 0 0 0
801 805Row Crop Agricult.   1 1 91 0 0 0
902 905Wetland              1 1 91 0 0 0
END GEN-INFO
```

ATEMP-DAT

```
*** <PLS >      ELDAT      AIRTEMP
*** x - x      (ft)      (deg F)
101 905      0.      30.
END ATEMP-DAT
```

SNOW-FLAGS

```
*** <PLS >
*** x - x SNOP VKM
101 905 1 0
END SNOW-FLAGS
```

SNOW-PARM1

```
*** < PLS>      LAT      MELEV      SHADE      SNOWCF      COVIND      KMELT
TBASE
*** x - x degrees (ft) (in) (in/d.F)
(F)
101 125      38.5      1672.      0.8      1.2      1.      0.12
32.
201 905      38.5      1672.      0.2      1.2      1.      0.12
32.
END SNOW-PARM1
```

SNOW-PARM2

```
*** <PLS >      RDCSN      TSNOW      SNOEVP      CCFACT      MWATER      MGMELT
*** x - x      (deg F)
101 905      0.15      32.      0.1      1.      0.03      0.01
END SNOW-PARM2
```

PWAT-PARM1

```
*** <PLS >
*** x - x CSNO RTOP UZFG VCS VUZ VNN VIFW VIRC VLE IFFC HWT IRRG
          Flags
101 905 1 1 1 1 0 0 0 0 1 1 0 0
END PWAT-PARM1
```

PWAT-PARM2

```
*** < PLS>      FOREST      LZSN      INFILT      LSUR      SLSUR      KVARV
AGWRC
*** x - x      (in) (in/hr) (ft) (1/in)
(1/day)
101 105      0.07      6.      0.012      100.      0.38      0.
0.9
111 115      0.07      6.      0.012      200.      0.2      0.
0.91
```

121	125	0.07	6.	0.012	300.	0.09	0.
0.93							
201	205	0.	5.	0.01	200.	0.16	0.
0.89							
301	305	0.	5.	0.005	200.	0.14	0.
0.89							
401	405	0.	8.	0.03	200.	0.33	0.
0.89							
601	605	0.	5.	0.015	200.	0.23	0.
0.89							
701	705	0.	5.	0.001	200.	0.11	0.
0.89							
801	805	0.	5.	0.015	200.	0.1	0.
0.89							
902	905	0.	2.	0.1	200.	0.09	0.
0.89							

END PWAT-PARM2

PWAT-PARM3

*** < PLS>	PETMAX	PETMIN	INFEXP	INFILD	DEEPPFR	BASETTP
AGWETP						
*** x - x	(deg F)	(deg F)				
101 111	39.	33.	2.	2.	0.	0.02
0.01						
121	39.	33.	2.	2.	0.	0.02
0.01						
201 905	39.	33.	2.	2.	0.	0.02
0.01						

END PWAT-PARM3

PWAT-PARM4

*** <PLS >	CEPSC	UZSN	NSUR	INTFW	IRC	LZETP
*** x - x	(in)	(in)			(1/day)	
101 105	0.14	0.7	0.35	3.	0.5	0.68
111 115	0.14	0.9	0.35	2.5	0.5	0.68
121 125	0.14	1.1	0.35	2.	0.5	0.68
201 205	0.05	0.5	0.7	1.5	0.5	0.46
301 305	0.01	0.3	0.1	1.	0.5	0.2
401 405	0.01	0.8	0.15	2.5	0.5	0.2
601 605	0.01	0.5	0.2	1.5	0.5	0.38
701 705	0.01	1.	0.01	1.	0.5	0.1
801 805	0.05	0.5	0.15	1.5	0.5	0.46
902 905	0.1	1.	0.5	1.	0.5	0.7

END PWAT-PARM4

PWAT-STATE1

*** < PLS>	PWATER state variables (in)						
*** x - x	CEPS	SURS	UZS	IFWS	LZS	AGWS	
GWVS							
101 905	0.01	0.01	0.7	0.01	5.	1.	
0.							

END PWAT-STATE1

MON-INTERCEP

*** <PLS >	Interception storage capacity at start of each month (in)											
*** x - x	JAN	FEB	MAR	APR	MAY	JUN	JUL	AUG	SEP	OCT	NOV	DEC
101 125	0.03	0.04	0.05	0.06	0.1	0.23	0.75	0.8	0.8	0.75	0.2	0.06
201 205	0.06	0.06	0.06	0.07	0.095	0.1	0.11	0.11	0.11	0.075	0.07	0.065
301 605	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01
701 705	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.	0.
801 805	0.03	0.03	0.03	0.03	0.01	0.05	0.12	0.14	0.14	0.01	0.05	0.04
902 905	0.03	0.03	0.03	0.03	0.04	0.05	0.06	0.06	0.06	0.05	0.04	0.03

END MON-INTERCEP

```

MON-LZETPARM
*** <PLS > Lower zone evapotransp parm at start of each month
*** x - x JAN FEB MAR APR MAY JUN JUL AUG SEP OCT NOV DEC
101 125 0.2 0.25 0.3 0.4 0.6 0.7 0.8 0.8 0.7 0.45 0.3 0.2
201 205 0.15 0.15 0.2 0.25 0.4 0.5 0.6 0.6 0.5 0.25 0.15 0.15
301 705 0.1 0.1 0.1 0.1 0.1 0.1 0.1 0.1 0.1 0.1 0.1 0.1 0.1
801 805 0.1 0.1 0.1 0.1 0.25 0.55 0.55 0.55 0.45 0.25 0.15 0.1
902 905 0.3 0.3 0.35 0.6 0.7 0.8 0.9 0.9 0.8 0.6 0.4 0.3
END MON-LZETPARM

END PERLND

RCHRES
ACTIVITY
*** RCHRES Active sections
*** x - x HYFG ADFG CNFG HTFG SDFG GQFG OXFG NUFG PKFG PHFG
1 27 1 0 0 0 0 0 0 0 0 0 0 0
END ACTIVITY

PRINT-INFO
*** RCHRES Printout level flags
*** x - x HYDR ADCA CONS HEAT SED GQL OXRX NUTR PLNK PHCB PIVL PYR
1 27 5 6 6 6 6 6 6 6 6 6 1 9
END PRINT-INFO

GEN-INFO
*** Name Nexits Unit Systems Printer
*** RCHRES t-series Engl Metr LKFG
*** x - x in out
1 27 1 1 1 0 0 0 0 0
END GEN-INFO

HYDR-PARM1
*** Flags for HYDR section
***RC HRES VC A1 A2 A3 ODFVFG for each *** ODGTFG for each FUNCT for
each
*** x - x FG FG FG FG possible exit *** possible exit possible
exit
1 27 0 1 1 1 4 0 0 0 0 0 0 0 0 0 1 1 1 1
1
END HYDR-PARM1

HYDR-PARM2
*** RCHRES FTBW FTBU LEN DELTH STCOR KS DB50
*** x - x (miles) (ft) (ft)
1 0. 1. 8.39 105. 3.2 0.5 0.01
2 0. 2. 6.93 154. 3.2 0.5 0.01
3 0. 3. 1.37 59. 3.2 0.5 0.01
*** BUFFALO CREEK
4 0. 4. 5.37 89. 3.2 0.5 0.01
5 0. 5. 4.73 108. 3.2 0.5 0.01
6 0. 6. 0.89 36. 3.2 0.5 0.01
7 0. 7. 8.8 72. 3.2 0.5 0.01
8 0. 8. 2.94 56. 3.2 0.5 0.01
*** TYGART AT ELKINS
9 0. 9. 4.25 62. 3.2 0.5 0.01
10 0. 10. 5.34 62. 3.2 0.5 0.01
11 0. 11. 2.19 82. 3.2 0.5 0.01
12 0. 12. 12.22 102. 3.2 0.5 0.01
13 0. 13. 3.21 112. 3.2 0.5 0.01
14 0. 14. 2.88 102. 3.2 0.5 0.01
15 0. 15. 20.63 617. 3.2 0.5 0.01

```

```

***          CLEAR FORK
  16      0.  16.      0.8      56.      3.2      0.5      0.01
  17      0.  17.      6.85     315.     3.2      0.5      0.01
  18      0.  18.      2.49     115.     3.2      0.5      0.01
  19      0.  19.      18.23    538.     3.2      0.5      0.01
  20      0.  20.      11.3     144.     3.2      0.5      0.01
***          BIG SANDY
  21      0.  21.      6.27     430.     3.2      0.5      0.01
  22      0.  22.      6.95      85.     3.2      0.5      0.01
  23      0.  23.      7.03     112.     3.2      0.5      0.01
  24      0.  24.      9.48     400.     3.2      0.5      0.01
  25      0.  25.      2.41     167.     3.2      0.5      0.01
  26      0.  26.      0.84      36.     3.2      0.5      0.01
  27      0.  27.      1.25     121.     3.2      0.5      0.01
END HYDR-PARM2

```

```

HYDR-INIT
***          Initial conditions for HYDR section
***RC HRES          VOL  CAT Initial value of COLIND          initial value of
OUTDGT
*** x  - x          ac-ft          for each possible exit  for each possible exit,ft3
  1  27          0.01          4.2  4.5  4.5  4.5  4.2          2.1  1.2  0.5  1.2
1.8
END HYDR-INIT

```

END RCHRES

FTABLES

```

FTABLE          3
rows cols          ***
  8      4
depth          area          volume  outflow1 ***
  0.          26.75          0.          0.
  0.18        27.04          4.72         3.39
  1.76        29.7          49.57        156.57
  2.2         34.12          62.78        226.99
  2.74        91.68          112.59       293.98
  3.29        93.52          163.42       539.09
  56.54       272.34          9903.        237376.95
  109.78      451.16          29163.131017580.19
END FTABLE 3

```

```

FTABLE          2
rows cols          ***
  8      4
depth          area          volume  outflow1 ***
  0.          2.36          0.          0.
  0.1         2.4          0.25         0.9
  1.05        2.71          2.66         41.33
  1.31        3.23          3.38         59.96
  1.64        8.43          6.12         79.
  1.97        8.65          8.92        145.43
  33.79       29.75          620.02       72022.73
  65.62       50.85          1902.67      324306.72
END FTABLE 2

```

```

FTABLE          1
rows cols          ***
  8      4
depth          area          volume  outflow1 ***
  0.          61.02          0.          0.
  0.26        61.55          16.2         9.47

```

2.64	66.39	168.41	437.37
3.3	74.45	212.73	633.94
4.13	203.87	379.77	812.22
4.96	207.23	549.59	1485.91
85.09	533.04	30210.79	604843.81
165.23	858.84	85981.12492556	25

END FTABLE 1

FTABLE 5

rows cols ***
8 4

depth	area	volume	outflow1	***
0.	2.93	0.	0.	
0.16	2.96	0.47	3.3	
1.59	3.27	4.92	152.46	
1.98	3.78	6.24	221.06	
2.48	10.11	11.2	287.17	
2.98	10.33	16.27	526.95	
51.09	31.04	1011.4	237053.84	
99.2	51.76	3003.21026153	75	

END FTABLE 5

FTABLE 6

rows cols ***
8 4

depth	area	volume	outflow1	***
0.	42.54	0.	0.	
0.41	42.83	17.64	48.84	
4.13	45.48	181.87	2256.04	
5.17	49.9	229.24	3269.72	
6.46	139.03	407.59	4149.57	
7.75	140.87	588.32	7576.08	
133.01	319.39	29415.452869977	25	
258.28	497.91	80605.34	11338560.	

END FTABLE 6

FTABLE 4

rows cols ***
8 4

depth	area	volume	outflow1	***
0.	29.25	0.	0.	
0.24	29.52	6.99	9.13	
2.38	31.98	72.83	421.17	
2.97	36.07	92.05	610.49	
3.72	98.32	164.51	784.19	
4.46	100.02	238.23	1435.42	
76.57	265.45	13415.94	595538.19	
148.69	430.88	38523.172479231	25	

END FTABLE 4

FTABLE 7

rows cols ***
8 4

depth	area	volume	outflow1	***
0.	41.61	0.	0.	
0.28	41.97	11.5	12.38	
2.75	45.19	119.48	571.5	
3.44	50.57	150.89	828.34	
4.3	138.72	269.27	1060.28	
5.16	140.96	389.57	1939.33	
88.61	358.32	21221.36	783841.31	
172.06	575.67	60190.883217712	75	

END FTABLE 7

```

FTABLE      8
rows cols   ***
 8      4
  depth      area      volume  outflow1 ***
  0.      115.93      0.      0.
  0.39    116.75      45.22    26.38
  3.89    124.22      466.73   1218.51
  4.86    136.65      588.45   1766.01
  6.07    379.91     1046.79  2243.88
  7.29    385.09     1511.42  4097.78
 125.12   887.77    76500.761566379.25
 242.95  1390.45  210720.06  6223554.
END FTABLE  8

```

```

FTABLE      11
rows cols   ***
 8      4
  depth      area      volume  outflow1 ***
  0.      122.09      0.      0.
  0.56    122.81      68.23    99.69
  5.57    129.3      700.37   4607.33
  6.97    140.11     881.74   6677.47
  8.71    394.2      1564.23  8429.71
 10.45    398.71     2254.56  15373.69
 179.35   835.76  106507.48  5592222.
 348.25  1272.82  284580.31 21502960.
END FTABLE  11

```

```

FTABLE      13
rows cols   ***
 8      4
  depth      area      volume  outflow1 ***
  0.      245.58      0.      0.
  0.51    247.1      126.09    64.26
  5.12    260.75     1295.82  2969.25
  6.4     283.5      1631.91  4303.36
  8.      795.53     2896.78  5440.34
  9.6     805.01     4176.82  9924.75
 164.75  1724.57  200411.5  3649837.
 319.9   2644.13  539318.13 14140778.
END FTABLE  13

```

```

FTABLE      14
rows cols   ***
 8      4
  depth      area      volume  outflow1 ***
  0.      15.32      0.      0.
  0.22    15.47      3.32     8.29
  2.16    16.83      34.7     382.43
  2.7     19.09      43.88    554.36
  3.37    51.82      78.51    713.88
  4.05    52.76     113.78   1307.44
  69.47   144.19    6556.01  552488.75
 134.89  235.61   18979.622321737.25
END FTABLE  14

```

```

FTABLE      15
rows cols   ***
 8      4
  depth      area      volume  outflow1 ***
  0.      320.65      0.      0.
  0.43    322.82     139.27   71.14

```

4.33	342.31	1434.85	3286.43
5.41	374.79	1808.21	4763.04
6.76	1045.87	3213.8	6039.48
8.12	1059.4	4637.7	11024.61
139.33	2372.15	229766.9241	48549.75
270.54	3684.89	627142.81	16319877.

END FTABLE 15

FTABLE 12

rows cols				***
8	4			
depth	area	volume	outflow1	***
0.	16.95	0.	0.	
0.21	17.12	3.66	8.11	
2.15	18.62	38.22	374.45	
2.69	21.13	48.34	542.8	
3.36	57.33	86.49	699.07	
4.03	58.38	125.35	1280.34	
69.17	159.73	7229.28	541442.69	
134.32	261.08	20935.61	2276182.	

END FTABLE 12

FTABLE 10

rows cols				***
8	4			
depth	area	volume	outflow1	***
0.	11.45	0.	0.	
0.21	11.56	2.46	8.24	
2.13	12.58	25.64	380.13	
2.67	14.28	32.43	551.04	
3.33	38.74	58.03	709.8	
4.	39.45	84.1	1300.05	
68.7	108.16	4858.9	550525.19	
133.39	176.88	14079.432315	942.75	

END FTABLE 10

FTABLE 9

rows cols				***
8	4			
depth	area	volume	outflow1	***
0.	104.84	0.	0.	
0.59	105.44	61.56	131.01	
5.86	110.87	631.48	6055.06	
7.32	119.91	794.87	8775.7	
9.15	337.88	1409.65	11069.75	
10.98	341.64	2031.32	20185.21	
188.46	707.2	95107.84	7296876.	
365.95	1072.76	253065.2	27934868.	

END FTABLE 9

FTABLE 16

rows cols				***
8	4			
depth	area	volume	outflow1	***
0.	2.38	0.	0.	
0.15	2.41	0.35	3.45	
1.48	2.66	3.73	159.19	
1.85	3.1	4.73	230.84	
2.31	8.25	8.5	300.52	
2.78	8.43	12.36	551.72	
47.65	25.9	782.51	252006.27	
92.52	43.38	2336.75	1098215.	

END FTABLE 16


```

FTABLE      17
rows cols   ***
 8      4
  depth      area      volume  outflow1 ***
  0.         38.15      0.         0.
  0.22       38.52      8.52       10.4
  2.22       41.85      88.91      479.96
  2.78       47.39      112.42     695.73
  3.47       128.77     201.07     895.23
  4.17       131.08     291.32     1639.29
  71.55      355.02     16667.86  688793.06
  138.93     578.95     48133.37  2886226.75
END FTABLE 17

```

```

FTABLE      18
rows cols   ***
 8      4
  depth      area      volume  outflow1 ***
  0.         18.03      0.         0.
  0.26       18.19      4.78       18.1
  2.64       19.62      49.69      835.35
  3.3        22.01      62.76      1210.8
  4.12       60.26      112.05     1551.38
  4.95       61.26      162.15     2838.2
  84.94      157.65     8917.23  1155685.88
  164.92     254.05     25382.83  4763437.
END FTABLE 18

```

```

FTABLE      19
rows cols   ***
 8      4
  depth      area      volume  outflow1 ***
  0.         170.84     0.         0.
  0.31       172.22     53.44      24.64
  3.12       184.61     553.68     1137.87
  3.89       205.26     698.8      1649.2
  4.87       565.87     1245.51    2105.1
  5.84       574.48     1800.61    3848.08
  100.28     1409.06    95457.15  1523058.5
  194.71     2243.64    267926.5  6178898.5
END FTABLE 19

```

```

FTABLE      20
rows cols   ***
 8      4
  depth      area      volume  outflow1 ***
  0.         173.92     0.         0.
  0.43       175.1      75.06      45.59
  4.3        185.7      773.43     2106.04
  5.38       203.38     974.71     3052.3
  6.72       567.42     1732.47    3870.72
  8.07       574.78     2500.13    7065.88
  138.45     1289.16    124015.96  2661241.5
  268.84     2003.54    338677.    10474989.
END FTABLE 20

```

```

FTABLE      22
rows cols   ***
 8      4
  depth      area      volume  outflow1 ***
  0.         37.21      0.         0.
  0.23       37.56      8.66       14.5

```

2.32	40.73	90.26	669.25
2.9	46.01	114.1	970.09
3.62	125.27	203.98	1246.95
4.34	127.47	295.46	2282.83
74.56	341.01	16742.31	951819.56
144.77	554.54	48182.423972628.25	

END FTABLE 22

FTABLE 21

rows cols ***
8 4

depth	area	volume	outflow1	***
0.	62.88	0.	0.	
0.3	63.4	19.23	14.76	
3.05	68.01	199.3	681.38	
3.81	75.71	251.56	987.57	
4.76	208.53	448.48	1261.2	
5.71	211.73	648.44	2305.7	
98.02	522.72	34545.57	916033.31	
190.32	833.7	97148.513724290.75		

END FTABLE 21

FTABLE 23

rows cols ***
8 4

depth	area	volume	outflow1	***
0.	100.44	0.	0.	
0.41	101.14	41.31	43.5	
4.1	107.42	426.02	2009.29	
5.12	117.9	537.	2912.09	
6.4	328.38	954.84	3696.28	
7.69	332.75	1378.27	6748.72	
131.94	756.11	69024.21	2559600.	
256.19	1179.48	189273.97	10119986.	

END FTABLE 23

FTABLE 25

rows cols ***
8 4

depth	area	volume	outflow1	***
0.	15.15	0.	0.	
0.48	15.24	7.24	116.1	
4.77	16.12	74.52	5364.22	
5.96	17.58	93.87	7774.4	
7.45	49.2	166.72	9840.63	
8.94	49.81	240.47	17956.77	
153.44	108.71	11694.2	6666999.5	
297.95	167.62	31659.68	25997188.	

END FTABLE 25

FTABLE 24

rows cols ***
8 4

depth	area	volume	outflow1	***
0.	12.76	0.	0.	
0.22	12.88	2.76	11.48	
2.15	14.01	28.78	529.96	
2.69	15.9	36.4	768.22	
3.36	43.14	65.13	989.37	
4.03	43.93	94.38	1812.01	
69.21	120.18	5442.84	766210.81	
134.4	196.43	15761.533220936.75		

END FTABLE 24

```

FTABLE      26
rows cols      ***
 8      4
  depth      area      volume      outflow1 ***
  0.         83.47      0.         0.
  0.3        84.16      25.08      25.91
  2.99       90.35      260.06     1196.13
  3.74       100.66     328.29     1733.66
  4.68       277.05     585.36     2214.88
  5.61       281.35     846.44     4049.52
  96.32      698.1      45267.19  1613566.88
 187.02     1114.86   127489.54  6571224.
END FTABLE 26

```

```

FTABLE      27
rows cols      ***
 8      4
  depth      area      volume      outflow1 ***
  0.         25.14      0.         0.
  0.51       25.29      12.89      218.24
  5.11       26.69      132.52     10084.57
  6.39       29.02      166.89     14615.63
  7.99       81.44      296.25     18477.5
  9.59       82.41      427.16     33708.39
 164.59     176.59     20500.     12398078.
 319.6      270.77    55171.63  48039260.
END FTABLE 27
END FTABLES

```

```

COPY
TIMESERIES
Copy-opn***
*** x - x NPT NMN
 1  5  0  7
END TIMESERIES

```

END COPY

```

EXT SOURCES
<-Volume-> <Member> SsysSgap<--Mult-->Tran <-Target vols> <-Grp> <-Member-> ***
<Name> x <Name> x tem strg<-factor->strg <Name> x x <Name> x x ***
*** Met Seg TWELVEPO
WDM 91 PREC ENGLZERO SAME PERLND 101 EXTNL PREC
WDM 103 ATEM ENGL SAME PERLND 101 EXTNL GATMP
WDM 106 PEVT ENGL SAME PERLND 101 EXTNL PETINP
*** Met Seg TWELVEPO
WDM 91 PREC ENGLZERO SAME PERLND 111 EXTNL PREC
WDM 103 ATEM ENGL SAME PERLND 111 EXTNL GATMP
WDM 106 PEVT ENGL SAME PERLND 111 EXTNL PETINP
*** Met Seg TWELVEPO
WDM 91 PREC ENGLZERO SAME PERLND 121 EXTNL PREC
WDM 103 ATEM ENGL SAME PERLND 121 EXTNL GATMP
WDM 106 PEVT ENGL SAME PERLND 121 EXTNL PETINP
*** Met Seg TWELVEPO
WDM 91 PREC ENGLZERO SAME PERLND 201 EXTNL PREC
WDM 103 ATEM ENGL SAME PERLND 201 EXTNL GATMP
WDM 106 PEVT ENGL SAME PERLND 201 EXTNL PETINP
*** Met Seg TWELVEPO
WDM 91 PREC ENGLZERO SAME PERLND 301 EXTNL PREC
WDM 103 ATEM ENGL SAME PERLND 301 EXTNL GATMP
WDM 106 PEVT ENGL SAME PERLND 301 EXTNL PETINP
*** Met Seg TWELVEPO

```

WDM	91	PREC	ENGLZERO	SAME	PERLND	401	EXTNL	PREC
WDM	103	ATEM	ENGL	SAME	PERLND	401	EXTNL	GATMP
WDM	106	PEVT	ENGL	SAME	PERLND	401	EXTNL	PETINP
*** Met Seg TWELVEPO								
WDM	91	PREC	ENGLZERO	SAME	PERLND	601	EXTNL	PREC
WDM	103	ATEM	ENGL	SAME	PERLND	601	EXTNL	GATMP
WDM	106	PEVT	ENGL	SAME	PERLND	601	EXTNL	PETINP
*** Met Seg TWELVEPO								
WDM	91	PREC	ENGLZERO	SAME	PERLND	701	EXTNL	PREC
WDM	103	ATEM	ENGL	SAME	PERLND	701	EXTNL	GATMP
WDM	106	PEVT	ENGL	SAME	PERLND	701	EXTNL	PETINP
*** Met Seg TWELVEPO								
WDM	91	PREC	ENGLZERO	SAME	PERLND	801	EXTNL	PREC
WDM	103	ATEM	ENGL	SAME	PERLND	801	EXTNL	GATMP
WDM	106	PEVT	ENGL	SAME	PERLND	801	EXTNL	PETINP
*** Met Seg BUFF								
WDM	211	PREC	ENGLZERO	SAME	PERLND	102	EXTNL	PREC
WDM	203	ATEM	ENGL	SAME	PERLND	102	EXTNL	GATMP
WDM	206	PEVT	ENGL	SAME	PERLND	102	EXTNL	PETINP
*** Met Seg BUFF								
WDM	211	PREC	ENGLZERO	SAME	PERLND	112	EXTNL	PREC
WDM	203	ATEM	ENGL	SAME	PERLND	112	EXTNL	GATMP
WDM	206	PEVT	ENGL	SAME	PERLND	112	EXTNL	PETINP
*** Met Seg BUFF								
WDM	211	PREC	ENGLZERO	SAME	PERLND	122	EXTNL	PREC
WDM	203	ATEM	ENGL	SAME	PERLND	122	EXTNL	GATMP
WDM	206	PEVT	ENGL	SAME	PERLND	122	EXTNL	PETINP
*** Met Seg BUFF								
WDM	211	PREC	ENGLZERO	SAME	PERLND	202	EXTNL	PREC
WDM	203	ATEM	ENGL	SAME	PERLND	202	EXTNL	GATMP
WDM	206	PEVT	ENGL	SAME	PERLND	202	EXTNL	PETINP
*** Met Seg BUFF								
WDM	211	PREC	ENGLZERO	SAME	PERLND	302	EXTNL	PREC
WDM	203	ATEM	ENGL	SAME	PERLND	302	EXTNL	GATMP
WDM	206	PEVT	ENGL	SAME	PERLND	302	EXTNL	PETINP
*** Met Seg BUFF								
WDM	211	PREC	ENGLZERO	SAME	PERLND	402	EXTNL	PREC
WDM	203	ATEM	ENGL	SAME	PERLND	402	EXTNL	GATMP
WDM	206	PEVT	ENGL	SAME	PERLND	402	EXTNL	PETINP
*** Met Seg BUFF								
WDM	211	PREC	ENGLZERO	SAME	PERLND	602	EXTNL	PREC
WDM	203	ATEM	ENGL	SAME	PERLND	602	EXTNL	GATMP
WDM	206	PEVT	ENGL	SAME	PERLND	602	EXTNL	PETINP
*** Met Seg BUFF								
WDM	211	PREC	ENGLZERO	SAME	PERLND	702	EXTNL	PREC
WDM	203	ATEM	ENGL	SAME	PERLND	702	EXTNL	GATMP
WDM	206	PEVT	ENGL	SAME	PERLND	702	EXTNL	PETINP
*** Met Seg BUFF								
WDM	211	PREC	ENGLZERO	SAME	PERLND	802	EXTNL	PREC
WDM	203	ATEM	ENGL	SAME	PERLND	802	EXTNL	GATMP
WDM	206	PEVT	ENGL	SAME	PERLND	802	EXTNL	PETINP
*** Met Seg BUFF								
WDM	211	PREC	ENGLZERO	SAME	PERLND	902	EXTNL	PREC
WDM	203	ATEM	ENGL	SAME	PERLND	902	EXTNL	GATMP
WDM	206	PEVT	ENGL	SAME	PERLND	902	EXTNL	PETINP
*** Met Seg TV ELKIN								
WDM	311	PREC	ENGLZERO	SAME	PERLND	103	EXTNL	PREC
WDM	303	ATEM	ENGL	SAME	PERLND	103	EXTNL	GATMP
WDM	306	PEVT	ENGL	SAME	PERLND	103	EXTNL	PETINP
*** Met Seg TV ELKIN								
WDM	311	PREC	ENGLZERO	SAME	PERLND	113	EXTNL	PREC
WDM	303	ATEM	ENGL	SAME	PERLND	113	EXTNL	GATMP
WDM	306	PEVT	ENGL	SAME	PERLND	113	EXTNL	PETINP

```

*** Met Seg TV ELKIN
WDM 311 PREC ENGLZERO SAME PERLND 123 EXTNL PREC
WDM 303 ATEM ENGL SAME PERLND 123 EXTNL GATMP
WDM 306 PEVT ENGL SAME PERLND 123 EXTNL PETINP
*** Met Seg TV ELKIN
WDM 311 PREC ENGLZERO SAME PERLND 203 EXTNL PREC
WDM 303 ATEM ENGL SAME PERLND 203 EXTNL GATMP
WDM 306 PEVT ENGL SAME PERLND 203 EXTNL PETINP
*** Met Seg TV ELKIN
WDM 311 PREC ENGLZERO SAME PERLND 303 EXTNL PREC
WDM 303 ATEM ENGL SAME PERLND 303 EXTNL GATMP
WDM 306 PEVT ENGL SAME PERLND 303 EXTNL PETINP
*** Met Seg TV ELKIN
WDM 311 PREC ENGLZERO SAME PERLND 403 EXTNL PREC
WDM 303 ATEM ENGL SAME PERLND 403 EXTNL GATMP
WDM 306 PEVT ENGL SAME PERLND 403 EXTNL PETINP
*** Met Seg TV ELKIN
WDM 311 PREC ENGLZERO SAME PERLND 603 EXTNL PREC
WDM 303 ATEM ENGL SAME PERLND 603 EXTNL GATMP
WDM 306 PEVT ENGL SAME PERLND 603 EXTNL PETINP
*** Met Seg TV ELKIN
WDM 311 PREC ENGLZERO SAME PERLND 703 EXTNL PREC
WDM 303 ATEM ENGL SAME PERLND 703 EXTNL GATMP
WDM 306 PEVT ENGL SAME PERLND 703 EXTNL PETINP
*** Met Seg TV ELKIN
WDM 311 PREC ENGLZERO SAME PERLND 803 EXTNL PREC
WDM 303 ATEM ENGL SAME PERLND 803 EXTNL GATMP
WDM 306 PEVT ENGL SAME PERLND 803 EXTNL PETINP
*** Met Seg TV ELKIN
WDM 311 PREC ENGLZERO SAME PERLND 903 EXTNL PREC
WDM 303 ATEM ENGL SAME PERLND 903 EXTNL GATMP
WDM 306 PEVT ENGL SAME PERLND 903 EXTNL PETINP
*** Met Seg CLEAR FO
WDM 411 PREC ENGLZERO SAME PERLND 104 EXTNL PREC
WDM 403 ATEM ENGL SAME PERLND 104 EXTNL GATMP
WDM 406 PEVT ENGL SAME PERLND 104 EXTNL PETINP
*** Met Seg CLEAR FO
WDM 411 PREC ENGLZERO SAME PERLND 114 EXTNL PREC
WDM 403 ATEM ENGL SAME PERLND 114 EXTNL GATMP
WDM 406 PEVT ENGL SAME PERLND 114 EXTNL PETINP
*** Met Seg CLEAR FO
WDM 411 PREC ENGLZERO SAME PERLND 124 EXTNL PREC
WDM 403 ATEM ENGL SAME PERLND 124 EXTNL GATMP
WDM 406 PEVT ENGL SAME PERLND 124 EXTNL PETINP
*** Met Seg CLEAR FO
WDM 411 PREC ENGLZERO SAME PERLND 204 EXTNL PREC
WDM 403 ATEM ENGL SAME PERLND 204 EXTNL GATMP
WDM 406 PEVT ENGL SAME PERLND 204 EXTNL PETINP
*** Met Seg CLEAR FO
WDM 411 PREC ENGLZERO SAME PERLND 304 EXTNL PREC
WDM 403 ATEM ENGL SAME PERLND 304 EXTNL GATMP
WDM 406 PEVT ENGL SAME PERLND 304 EXTNL PETINP
*** Met Seg CLEAR FO
WDM 411 PREC ENGLZERO SAME PERLND 404 EXTNL PREC
WDM 403 ATEM ENGL SAME PERLND 404 EXTNL GATMP
WDM 406 PEVT ENGL SAME PERLND 404 EXTNL PETINP
*** Met Seg CLEAR FO
WDM 411 PREC ENGLZERO SAME PERLND 604 EXTNL PREC
WDM 403 ATEM ENGL SAME PERLND 604 EXTNL GATMP
WDM 406 PEVT ENGL SAME PERLND 604 EXTNL PETINP
*** Met Seg CLEAR FO
WDM 411 PREC ENGLZERO SAME PERLND 704 EXTNL PREC
WDM 403 ATEM ENGL SAME PERLND 704 EXTNL GATMP

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WDM	406	PEVT	ENGL	SAME	PERLND	704	EXTNL	PETINP	
***	Met Seg CLEAR FO								
WDM	411	PREC	ENGLZERO	SAME	PERLND	804	EXTNL	PREC	
WDM	403	ATEM	ENGL	SAME	PERLND	804	EXTNL	GATMP	
WDM	406	PEVT	ENGL	SAME	PERLND	804	EXTNL	PETINP	
***	Met Seg CLEAR FO								
WDM	411	PREC	ENGLZERO	SAME	PERLND	904	EXTNL	PREC	
WDM	403	ATEM	ENGL	SAME	PERLND	904	EXTNL	GATMP	
WDM	406	PEVT	ENGL	SAME	PERLND	904	EXTNL	PETINP	
***	Met Seg BIG SAND,PI:GATMP=0.83,PI:PETINP=0.83								
WDM	511	PREC	ENGLZERO	SAME	PERLND	105	EXTNL	PREC	
WDM	503	ATEM	ENGL	0.83	SAME	PERLND	105	EXTNL	GATMP
WDM	506	PEVT	ENGL	0.83	SAME	PERLND	105	EXTNL	PETINP
***	Met Seg BIG SAND,PI:GATMP=0.83,PI:PETINP=0.83								
WDM	511	PREC	ENGLZERO	SAME	PERLND	115	EXTNL	PREC	
WDM	503	ATEM	ENGL	0.83	SAME	PERLND	115	EXTNL	GATMP
WDM	506	PEVT	ENGL	0.83	SAME	PERLND	115	EXTNL	PETINP
***	Met Seg BIG SAND,PI:GATMP=0.83,PI:PETINP=0.83								
WDM	511	PREC	ENGLZERO	SAME	PERLND	125	EXTNL	PREC	
WDM	503	ATEM	ENGL	0.83	SAME	PERLND	125	EXTNL	GATMP
WDM	506	PEVT	ENGL	0.83	SAME	PERLND	125	EXTNL	PETINP
***	Met Seg BIG SAND,PI:GATMP=0.83,PI:PETINP=0.83								
WDM	511	PREC	ENGLZERO	SAME	PERLND	205	EXTNL	PREC	
WDM	503	ATEM	ENGL	0.83	SAME	PERLND	205	EXTNL	GATMP
WDM	506	PEVT	ENGL	0.83	SAME	PERLND	205	EXTNL	PETINP
***	Met Seg BIG SAND,PI:GATMP=0.83,PI:PETINP=0.83								
WDM	511	PREC	ENGLZERO	SAME	PERLND	305	EXTNL	PREC	
WDM	503	ATEM	ENGL	0.83	SAME	PERLND	305	EXTNL	GATMP
WDM	506	PEVT	ENGL	0.83	SAME	PERLND	305	EXTNL	PETINP
***	Met Seg BIG SAND,PI:GATMP=0.83,PI:PETINP=0.83								
WDM	511	PREC	ENGLZERO	SAME	PERLND	405	EXTNL	PREC	
WDM	503	ATEM	ENGL	0.83	SAME	PERLND	405	EXTNL	GATMP
WDM	506	PEVT	ENGL	0.83	SAME	PERLND	405	EXTNL	PETINP
***	Met Seg BIG SAND,PI:GATMP=0.83,PI:PETINP=0.83								
WDM	511	PREC	ENGLZERO	SAME	PERLND	605	EXTNL	PREC	
WDM	503	ATEM	ENGL	0.83	SAME	PERLND	605	EXTNL	GATMP
WDM	506	PEVT	ENGL	0.83	SAME	PERLND	605	EXTNL	PETINP
***	Met Seg BIG SAND,PI:GATMP=0.83,PI:PETINP=0.83								
WDM	511	PREC	ENGLZERO	SAME	PERLND	705	EXTNL	PREC	
WDM	503	ATEM	ENGL	0.83	SAME	PERLND	705	EXTNL	GATMP
WDM	506	PEVT	ENGL	0.83	SAME	PERLND	705	EXTNL	PETINP
***	Met Seg BIG SAND,PI:GATMP=0.83,PI:PETINP=0.83								
WDM	511	PREC	ENGLZERO	SAME	PERLND	805	EXTNL	PREC	
WDM	503	ATEM	ENGL	0.83	SAME	PERLND	805	EXTNL	GATMP
WDM	506	PEVT	ENGL	0.83	SAME	PERLND	805	EXTNL	PETINP
***	Met Seg BIG SAND,PI:GATMP=0.83,PI:PETINP=0.83								
WDM	511	PREC	ENGLZERO	SAME	PERLND	905	EXTNL	PREC	
WDM	503	ATEM	ENGL	0.83	SAME	PERLND	905	EXTNL	GATMP
WDM	506	PEVT	ENGL	0.83	SAME	PERLND	905	EXTNL	PETINP
***	Met Seg TWELVEPO								
WDM	91	PREC	ENGLZERO	SAME	RCHRES	1	3	EXTNL	PREC
WDM	103	ATEM	ENGL	SAME	RCHRES	1	3	EXTNL	GATMP
***	Met Seg BUFF								
WDM	211	PREC	ENGLZERO	SAME	RCHRES	4	8	EXTNL	PREC
WDM	203	ATEM	ENGL	SAME	RCHRES	4	8	EXTNL	GATMP
***	Met Seg TV ELKIN								
WDM	311	PREC	ENGLZERO	SAME	RCHRES	9	15	EXTNL	PREC
WDM	303	ATEM	ENGL	SAME	RCHRES	9	15	EXTNL	GATMP
***	Met Seg CLEAR FO								
WDM	411	PREC	ENGLZERO	SAME	RCHRES	16	20	EXTNL	PREC
WDM	403	ATEM	ENGL	SAME	RCHRES	16	20	EXTNL	GATMP
***	Met Seg BIG SAND,PI:GATMP=0.83,PI:PETINP=0.83								
WDM	511	PREC	ENGLZERO	SAME	RCHRES	21	27	EXTNL	PREC

WDM 503 ATEM ENGL 0.83SAME RCHRES 21 27 EXTNL GATMP
 END EXT SOURCES

SCHEMATIC

<-Volume-> <Name> x	<--Area--> <-factor-->	<-Volume-> <Name> x	<ML#>	*** ***	<sb> x x
***	TWELVE POLE CREEK				
PERLND 301		RCHRES 3	2		
PERLND 601		RCHRES 3	2		
PERLND 401	330	RCHRES 3	2		
PERLND 121	203	RCHRES 3	2		
PERLND 111	324	RCHRES 3	2		
PERLND 101	1492	RCHRES 3	2		
PERLND 201	18	RCHRES 2	2		
PERLND 201	87	RCHRES 2	2		
PERLND 801	2	RCHRES 2	2		
PERLND 301	17	RCHRES 2	2		
PERLND 601	20	RCHRES 2	2		
PERLND 401	170	RCHRES 2	2		
PERLND 121	755	RCHRES 2	2		
PERLND 111	1133	RCHRES 2	2		
PERLND 101	6305	RCHRES 2	2		
PERLND 201	6	RCHRES 1	2		
PERLND 201	371	RCHRES 1	2		
PERLND 301	14	RCHRES 1	2		
PERLND 601	46	RCHRES 1	2		
PERLND 701	7	RCHRES 1	2		
PERLND 401	2154	RCHRES 1	2		
PERLND 121	1080	RCHRES 1	2		
PERLND 111	1829	RCHRES 1	2		
PERLND 101	7281	RCHRES 1	2		
RCHRES 3		RCHRES 1	3		
RCHRES 2		RCHRES 1	3		
***	BUFFALO CREEK				
PERLND 202	470	RCHRES 5	2		
PERLND 202	1529	RCHRES 5	2		
PERLND 302	117	RCHRES 5	2		
PERLND 602	40	RCHRES 5	2		
PERLND 902	25	RCHRES 5	2		
PERLND 702	50	RCHRES 5	2		
PERLND 402	2	RCHRES 5	2		
PERLND 122	2382	RCHRES 5	2		
PERLND 112	3884	RCHRES 5	2		
PERLND 102	9668	RCHRES 5	2		
PERLND 202	219	RCHRES 6	2		
PERLND 202	1752	RCHRES 6	2		
PERLND 302	38	RCHRES 6	2		
PERLND 602	3	RCHRES 6	2		
PERLND 902	15	RCHRES 6	2		
PERLND 702	3	RCHRES 6	2		
PERLND 122	1262	RCHRES 6	2		
PERLND 112	1678	RCHRES 6	2		
PERLND 102	1635	RCHRES 6	2		
PERLND 202	380	RCHRES 4	2		
PERLND 202	2679	RCHRES 4	2		
PERLND 302	229	RCHRES 4	2		
PERLND 602	9	RCHRES 4	2		
PERLND 902	23	RCHRES 4	2		
PERLND 702	111	RCHRES 4	2		
PERLND 402	659	RCHRES 4	2		
PERLND 122	3869	RCHRES 4	2		
PERLND 112	6121	RCHRES 4	2		
PERLND 102	12089	RCHRES 4	2		

PERLND 202	507	RCHRES	7	2
PERLND 202	3640	RCHRES	7	2
PERLND 302	802	RCHRES	7	2
PERLND 602	79	RCHRES	7	2
PERLND 902	14	RCHRES	7	2
PERLND 702	18	RCHRES	7	2
PERLND 402	197	RCHRES	7	2
PERLND 122	2725	RCHRES	7	2
PERLND 112	3658	RCHRES	7	2
PERLND 102	6036	RCHRES	7	2
RCHRES 5		RCHRES	7	3
RCHRES 4		RCHRES	7	3
PERLND 202	54	RCHRES	8	2
PERLND 202	872	RCHRES	8	2
PERLND 302	102	RCHRES	8	2
PERLND 602	2	RCHRES	8	2
PERLND 702		RCHRES	8	2
PERLND 402	26	RCHRES	8	2
PERLND 122	752	RCHRES	8	2
PERLND 112	939	RCHRES	8	2
PERLND 102	892	RCHRES	8	2
RCHRES 6		RCHRES	8	3
RCHRES 7		RCHRES	8	3

	TV ELK			
PERLND 203	37	RCHRES	11	2
PERLND 203	2749	RCHRES	11	2
PERLND 803	21	RCHRES	11	2
PERLND 303	410	RCHRES	11	2
PERLND 603	179	RCHRES	11	2
PERLND 903	44	RCHRES	11	2
PERLND 703	34	RCHRES	11	2
PERLND 403	301	RCHRES	11	2
PERLND 123	1462	RCHRES	11	2
PERLND 113	2125	RCHRES	11	2
PERLND 103	6487	RCHRES	11	2
PERLND 203	13	RCHRES	13	2
PERLND 203	1944	RCHRES	13	2
PERLND 103	3	RCHRES	13	2
PERLND 303	62	RCHRES	13	2
PERLND 603	33	RCHRES	13	2
PERLND 903	31	RCHRES	13	2
PERLND 703	30	RCHRES	13	2
PERLND 123	1636	RCHRES	13	2
PERLND 113	2630	RCHRES	13	2
PERLND 103	7710	RCHRES	13	2
PERLND 203	61	RCHRES	14	2
PERLND 203	792	RCHRES	14	2
PERLND 103	80	RCHRES	14	2
PERLND 303	159	RCHRES	14	2
PERLND 603	4	RCHRES	14	2
PERLND 903	42	RCHRES	14	2
PERLND 703	12	RCHRES	14	2
PERLND 123	3138	RCHRES	14	2
PERLND 113	3467	RCHRES	14	2
PERLND 103	6482	RCHRES	14	2
PERLND 203	1093	RCHRES	15	2
PERLND 203	9669	RCHRES	15	2
PERLND 103	126	RCHRES	15	2
PERLND 303	313	RCHRES	15	2
PERLND 603	74	RCHRES	15	2
PERLND 903	328	RCHRES	15	2
PERLND 703	531	RCHRES	15	2
PERLND 403	73	RCHRES	15	2

PERLND 123	10516	RCHRES 15	2
PERLND 113	13916	RCHRES 15	2
PERLND 103	44497	RCHRES 15	2
PERLND 203	225	RCHRES 12	2
PERLND 203	7159	RCHRES 12	2
PERLND 103	66	RCHRES 12	2
PERLND 303	232	RCHRES 12	2
PERLND 603	159	RCHRES 12	2
PERLND 903	231	RCHRES 12	2
PERLND 703	337	RCHRES 12	2
PERLND 123	3282	RCHRES 12	2
PERLND 113	4162	RCHRES 12	2
PERLND 103	12140	RCHRES 12	2
RCHRES 14		RCHRES 12	3
RCHRES 15		RCHRES 12	3
PERLND 203	77	RCHRES 10	2
PERLND 203	3130	RCHRES 10	2
PERLND 103	54	RCHRES 10	2
PERLND 303	371	RCHRES 10	2
PERLND 603	97	RCHRES 10	2
PERLND 903	378	RCHRES 10	2
PERLND 703	169	RCHRES 10	2
PERLND 123	1596	RCHRES 10	2
PERLND 113	1939	RCHRES 10	2
PERLND 103	7240	RCHRES 10	2
RCHRES 13		RCHRES 10	3
RCHRES 12		RCHRES 10	3
PERLND 203	53	RCHRES 9	2
PERLND 203	1396	RCHRES 9	2
PERLND 103	5	RCHRES 9	2
PERLND 303	1103	RCHRES 9	2
PERLND 603	80	RCHRES 9	2
PERLND 903	127	RCHRES 9	2
PERLND 703	193	RCHRES 9	2
PERLND 403	136	RCHRES 9	2
PERLND 123	724	RCHRES 9	2
PERLND 113	717	RCHRES 9	2
PERLND 103	1750	RCHRES 9	2
RCHRES 11		RCHRES 9	3
RCHRES 10		RCHRES 9	3

	CLEAR FORK		
PERLND 204	199	RCHRES 16	2
PERLND 204	25	RCHRES 16	2
PERLND 804	7	RCHRES 16	2
PERLND 304	46	RCHRES 16	2
PERLND 604	40	RCHRES 16	2
PERLND 704	3	RCHRES 16	2
PERLND 404	136	RCHRES 16	2
PERLND 124	170	RCHRES 16	2
PERLND 114	279	RCHRES 16	2
PERLND 104	4642	RCHRES 16	2
PERLND 204	109	RCHRES 17	2
PERLND 204	57	RCHRES 17	2
PERLND 804	22	RCHRES 17	2
PERLND 304	12	RCHRES 17	2
PERLND 604	47	RCHRES 17	2
PERLND 904	3	RCHRES 17	2
PERLND 704	218	RCHRES 17	2
PERLND 404	543	RCHRES 17	2
PERLND 124	664	RCHRES 17	2
PERLND 114	1128	RCHRES 17	2
PERLND 104	12525	RCHRES 17	2
PERLND 204	35	RCHRES 18	2

PERLND 204	30	RCHRES 18	2
PERLND 304	40	RCHRES 18	2
PERLND 604	12	RCHRES 18	2
PERLND 704	46	RCHRES 18	2
PERLND 404	673	RCHRES 18	2
PERLND 124	111	RCHRES 18	2
PERLND 114	133	RCHRES 18	2
PERLND 104	1576	RCHRES 18	2
RCHRES 16		RCHRES 18	3
RCHRES 17		RCHRES 18	3
PERLND 204	381	RCHRES 19	2
PERLND 204	574	RCHRES 19	2
PERLND 804	16	RCHRES 19	2
PERLND 304	559	RCHRES 19	2
PERLND 604	158	RCHRES 19	2
PERLND 904	7	RCHRES 19	2
PERLND 704	14	RCHRES 19	2
PERLND 404	1349	RCHRES 19	2
PERLND 124	2516	RCHRES 19	2
PERLND 114	4201	RCHRES 19	2
PERLND 104	25874	RCHRES 19	2
PERLND 204	423	RCHRES 20	2
PERLND 204	767	RCHRES 20	2
PERLND 304	430	RCHRES 20	2
PERLND 604	167	RCHRES 20	2
PERLND 904	47	RCHRES 20	2
PERLND 704	248	RCHRES 20	2
PERLND 404	964	RCHRES 20	2
PERLND 124	1341	RCHRES 20	2
PERLND 114	1563	RCHRES 20	2
PERLND 104	14732	RCHRES 20	2
RCHRES 18		RCHRES 20	3
RCHRES 19		RCHRES 20	3

	BIG SANDY		
PERLND 205	2394	RCHRES 22	2
PERLND 805	560	RCHRES 22	2
PERLND 305	28	RCHRES 22	2
PERLND 605	16	RCHRES 22	2
PERLND 905	8	RCHRES 22	2
PERLND 705	10	RCHRES 22	2
PERLND 405	23	RCHRES 22	2
PERLND 125	15780	RCHRES 22	2
PERLND 115	10285	RCHRES 22	2
PERLND 105	4568	RCHRES 22	2
PERLND 205	2339	RCHRES 21	2
PERLND 805	590	RCHRES 21	2
PERLND 305	12	RCHRES 21	2
PERLND 605		RCHRES 21	2
PERLND 905	415	RCHRES 21	2
PERLND 705	235	RCHRES 21	2
PERLND 405	18	RCHRES 21	2
PERLND 125	10497	RCHRES 21	2
PERLND 115	2399	RCHRES 21	2
PERLND 105	470	RCHRES 21	2
PERLND 205	3710	RCHRES 23	2
PERLND 805	2013	RCHRES 23	2
PERLND 305	66	RCHRES 23	2
PERLND 905	83	RCHRES 23	2
PERLND 705	154	RCHRES 23	2
PERLND 405	58	RCHRES 23	2
PERLND 125	9286	RCHRES 23	2
PERLND 115	4138	RCHRES 23	2
PERLND 105	609	RCHRES 23	2

RCHRES	22		RCHRES	23	3
RCHRES	21		RCHRES	23	3
PERLND	205	671	RCHRES	25	2
PERLND	805	389	RCHRES	25	2
PERLND	305	17	RCHRES	25	2
PERLND	905	29	RCHRES	25	2
PERLND	705	96	RCHRES	25	2
PERLND	405	17	RCHRES	25	2
PERLND	125	5641	RCHRES	25	2
PERLND	115	4874	RCHRES	25	2
PERLND	105	2379	RCHRES	25	2
PERLND	205	6728	RCHRES	24	2
PERLND	805	3399	RCHRES	24	2
PERLND	305	90	RCHRES	24	2
PERLND	905	231	RCHRES	24	2
PERLND	705	33	RCHRES	24	2
PERLND	405	186	RCHRES	24	2
PERLND	125	15322	RCHRES	24	2
PERLND	115	5163	RCHRES	24	2
PERLND	105	1080	RCHRES	24	2
PERLND	705	19	RCHRES	26	2
PERLND	125	55	RCHRES	26	2
PERLND	115	53	RCHRES	26	2
PERLND	105	105	RCHRES	26	2
RCHRES	23		RCHRES	26	3
RCHRES	24		RCHRES	26	3
PERLND	205	935	RCHRES	27	2
PERLND	805	649	RCHRES	27	2
PERLND	305	3	RCHRES	27	2
PERLND	905	16	RCHRES	27	2
PERLND	705	36	RCHRES	27	2
PERLND	405	38	RCHRES	27	2
PERLND	125	2030	RCHRES	27	2
PERLND	115	1454	RCHRES	27	2
PERLND	105	525	RCHRES	27	2
RCHRES	25		RCHRES	27	3
RCHRES	26		RCHRES	27	3
PERLND	201	24	COPY	1	90
PERLND	201	458	COPY	1	90
PERLND	301	32	COPY	1	90
PERLND	601	67	COPY	1	90
PERLND	701	7	COPY	1	90
PERLND	401	2654	COPY	1	90
PERLND	121	2038	COPY	1	90
PERLND	111	3286	COPY	1	90
PERLND	101	15078	COPY	1	90
PERLND	801	2	COPY	1	90
PERLND	202	1630	COPY	2	90
PERLND	202	10472	COPY	2	90
PERLND	302	1288	COPY	2	90
PERLND	602	133	COPY	2	90
PERLND	702	183	COPY	2	90
PERLND	402	884	COPY	2	90
PERLND	122	10990	COPY	2	90
PERLND	112	16280	COPY	2	90
PERLND	102	30320	COPY	2	90
PERLND	902	77	COPY	2	90
PERLND	203	1559	COPY	3	90
PERLND	203	26839	COPY	3	90
PERLND	103	86640	COPY	3	90
PERLND	303	2650	COPY	3	90
PERLND	603	626	COPY	3	90
PERLND	903	1181	COPY	3	90

PERLND 703	1306	COPY	3	90
PERLND 403	510	COPY	3	90
PERLND 123	22354	COPY	3	90
PERLND 113	28956	COPY	3	90
PERLND 803	21	COPY	3	90
PERLND 204	1147	COPY	4	90
PERLND 204	1453	COPY	4	90
PERLND 304	1087	COPY	4	90
PERLND 604	424	COPY	4	90
PERLND 904	57	COPY	4	90
PERLND 704	529	COPY	4	90
PERLND 404	3665	COPY	4	90
PERLND 124	4802	COPY	4	90
PERLND 114	7304	COPY	4	90
PERLND 104	59349	COPY	4	90
PERLND 804	45	COPY	4	90
PERLND 205	16777	COPY	5	90
PERLND 805	7600	COPY	5	90
PERLND 305	216	COPY	5	90
PERLND 905	782	COPY	5	90
PERLND 705	583	COPY	5	90
PERLND 405	340	COPY	5	90
PERLND 125	58611	COPY	5	90
PERLND 115	28366	COPY	5	90
PERLND 105	9736	COPY	5	90
PERLND 605	17	COPY	5	90

END SCHEMATIC

EXT TARGETS

<-Volume-> <-Grp> <-Member-><--Mult-->Tran <-Volume-> <Member> Tsys Aggr Amd

<Name>	x	<Name>	x	x<-factor-->	strg	<Name>	x	<Name>	qf	tem	strg
strg***											
RCHRES	1	HYDR	RO	1	1	AVER	WDM	175	FLOW	1	ENGL AGGR REPL
RCHRES	1	ROFLOW	ROVOL	1	1	5.0749e-4	WDM	1001	SIMQ	1	ENGL AGGR REPL
RCHRES	8	HYDR	RO	1	1	AVER	WDM	275	FLOW	1	ENGL AGGR REPL
RCHRES	8	ROFLOW	ROVOL	1	1	1.6607e-4	WDM	2001	SIMQ	1	ENGL AGGR REPL
RCHRES	9	HYDR	RO	1	1	AVER	WDM	375	FLOW	1	ENGL AGGR REPL
RCHRES	9	ROFLOW	ROVOL	1	1	6.9508e-5	WDM	3001	SIMQ	1	ENGL AGGR REPL
RCHRES	20	HYDR	RO	1	1	AVER	WDM	475	FLOW	1	ENGL AGGR REPL
RCHRES	20	ROFLOW	ROVOL	1	1	1.5026e-4	WDM	4001	SIMQ	1	ENGL AGGR REPL
RCHRES	27	HYDR	RO	1	1	AVER	WDM	575	FLOW	1	ENGL AGGR REPL
RCHRES	27	ROFLOW	ROVOL	1	1	9.7539e-5	WDM	5001	SIMQ	1	ENGL AGGR REPL
COPY	1	OUTPUT	MEAN	1	1	4.229E-05	WDM	1002	SURO	1	ENGL
COPY	1	OUTPUT	MEAN	2	1	4.229E-05	WDM	1003	IFWO	1	ENGL AGGR REPL
COPY	1	OUTPUT	MEAN	3	1	4.229E-05	WDM	1004	AGWO	1	ENGL AGGR REPL
COPY	1	OUTPUT	MEAN	4	1	4.229E-05	WDM	1005	PETX	1	ENGL AGGR REPL
COPY	1	OUTPUT	MEAN	5	1	4.229E-05	WDM	1006	SAET	1	ENGL AGGR REPL
COPY	1	OUTPUT	MEAN	6	1	4.229E-05	AVER WDM	1007	UZSX	1	ENGL AGGR REPL
COPY	1	OUTPUT	MEAN	7	1	4.229E-05	AVER WDM	1008	LZSX	1	ENGL AGGR REPL
COPY	2	OUTPUT	MEAN	1	1	1.3839e-5	WDM	2002	SURO	1	ENGL AGGR REPL
COPY	2	OUTPUT	MEAN	2	1	1.3839e-5	WDM	2003	IFWO	1	ENGL AGGR REPL
COPY	2	OUTPUT	MEAN	3	1	1.3839e-5	WDM	2004	AGWO	1	ENGL AGGR REPL
COPY	2	OUTPUT	MEAN	4	1	1.3839e-5	WDM	2005	PETX	1	ENGL AGGR REPL
COPY	2	OUTPUT	MEAN	5	1	1.3839e-5	WDM	2006	SAET	1	ENGL AGGR REPL
COPY	2	OUTPUT	MEAN	6	1	1.3839e-5	AVER WDM	2007	UZSX	1	ENGL AGGR REPL
COPY	2	OUTPUT	MEAN	7	1	1.3839e-5	AVER WDM	2008	LZSX	1	ENGL AGGR REPL
COPY	3	OUTPUT	MEAN	1	1	5.7923e-6	WDM	3002	SURO	1	ENGL AGGR REPL
COPY	3	OUTPUT	MEAN	2	1	5.7923e-6	WDM	3003	IFWO	1	ENGL AGGR REPL
COPY	3	OUTPUT	MEAN	3	1	5.7923e-6	WDM	3004	AGWO	1	ENGL AGGR REPL
COPY	3	OUTPUT	MEAN	4	1	5.7923e-6	WDM	3005	PETX	1	ENGL AGGR REPL
COPY	3	OUTPUT	MEAN	5	1	5.7923e-6	WDM	3006	SAET	1	ENGL AGGR REPL
COPY	3	OUTPUT	MEAN	6	1	5.7923e-6	AVER WDM	3007	UZSX	1	ENGL AGGR REPL

```

COPY      3 OUTPUT MEAN      7 1 5.7923e-6AVER WDM   3008 LZSX   1 ENGL AGGR REPL
COPY      4 OUTPUT MEAN      1 1 1.2522e-5      WDM   4002 SURO   1 ENGL AGGR REPL
COPY      4 OUTPUT MEAN      2 1 1.2522e-5      WDM   4003 IFWO   1 ENGL AGGR REPL
COPY      4 OUTPUT MEAN      3 1 1.2522e-5      WDM   4004 AGWO   1 ENGL AGGR REPL
COPY      4 OUTPUT MEAN      4 1 1.2522e-5      WDM   4005 PETX   1 ENGL AGGR REPL
COPY      4 OUTPUT MEAN      5 1 1.2522e-5      WDM   4006 SAET   1 ENGL AGGR REPL
COPY      4 OUTPUT MEAN      6 1 1.2522e-5AVER WDM   4007 UZSX   1 ENGL AGGR REPL
COPY      4 OUTPUT MEAN      7 1 1.2522e-5AVER WDM   4008 LZSX   1 ENGL AGGR REPL
COPY      5 OUTPUT MEAN      1 1 8.1282e-6      WDM   5002 SURO   1 ENGL AGGR REPL
COPY      5 OUTPUT MEAN      2 1 8.1282e-6      WDM   5003 IFWO   1 ENGL AGGR REPL
COPY      5 OUTPUT MEAN      3 1 8.1282e-6      WDM   5004 AGWO   1 ENGL AGGR REPL
COPY      5 OUTPUT MEAN      4 1 8.1282e-6      WDM   5005 PETX   1 ENGL AGGR REPL
COPY      5 OUTPUT MEAN      5 1 8.1282e-6      WDM   5006 SAET   1 ENGL AGGR REPL
COPY      5 OUTPUT MEAN      6 1 8.1282e-6AVER WDM   5007 UZSX   1 ENGL AGGR REPL
COPY      5 OUTPUT MEAN      7 1 8.1282e-6AVER WDM   5008 LZSX   1 ENGL AGGR REPL
END EXT TARGETS

```

MASS-LINK

```

      MASS-LINK          2
<-Volume-> <-Grp> <-Member-><--Mult--> <-Target vols> <-Grp> <-Member->
***
<Name>          <Name> x x<-factor->          <Name>          <Name> x x
***
PERLND      PWATER PERO          0.0833333      RCHRES          INFLOW IVOL
      END MASS-LINK          2

```

```

      MASS-LINK          3
<-Volume-> <-Grp> <-Member-><--Mult--> <-Target vols> <-Grp> <-Member->
***
<Name>          <Name> x x<-factor->          <Name>          <Name> x x
***
RCHRES      ROFLOW          RCHRES          INFLOW
      END MASS-LINK          3

```

```

      MASS-LINK          90
<-Volume-> <-Grp> <-Member-><--Mult--> <-Target vols> <-Grp> <-Member->
***
<Name>          <Name> x x<-factor->          <Name>          <Name> x x
***
PERLND      PWATER SURO          COPY          INPUT MEAN 1
PERLND      PWATER IFWO          COPY          INPUT MEAN 2
PERLND      PWATER AGWO          COPY          INPUT MEAN 3
PERLND      PWATER PET          COPY          INPUT MEAN 4
PERLND      PWATER TAET          COPY          INPUT MEAN 5
PERLND      PWATER UZS          COPY          INPUT MEAN 6
PERLND      PWATER LZS          COPY          INPUT MEAN 7
      END MASS-LINK          90
END MASS-LINK

```

END RUN

Appendix C. Cyclic and Asymptotic Methods User's Control Input File (UCI)

```

RUN

GLOBAL
  UCI Created by WinHSPF for Mine_UCI
  START      1975/01/01 01:00  END      1995/12/30 23:00
  RUN INTERP  OUTPT LEVELS    1      0
  RESUME     0 RUN      1
  UNITS      2
END GLOBAL

FILES
<FILE> <UN#>***<----FILE NAME----->
MESSU    24  Mine_Site__emp_mm.ech
WDM      26  MineSite.wdm
         30  D:\Matlab_Functions\samsfunctions\Output_emp.txt
END FILES

OPN SEQUENCE
  INGRP          INDELT 01:00
  PERLND         101
  PLTGEN         1
  END INGRP
END OPN SEQUENCE

PERLND
  ACTIVITY
  *** <PLS >          Active Sections          ***
  *** x - x  ATMP  SNOW  PWAT  SED  PST  PWG  PQAL  MSTL  PEST  NITR  PHOS  TRAC  ***
  101      0    0    1    0    0    0    0    0    0    0    0    0    0
  END ACTIVITY

  PRINT-INFO
  *** < PLS>          Print-flags          PIVL  PYR
  *** x - x  ATMP  SNOW  PWAT  SED  PST  PWG  PQAL  MSTL  PEST  NITR  PHOS  TRAC
  101      6    6    6    6    6    6    6    6    6    6    6    6    6    1    12
  END PRINT-INFO

  GEN-INFO
  ***          Name          Unit-systems  Printer  BinaryOut
  *** <PLS >          t-series  Engl  Metr  Engl  Metr
  *** x - x          in  out
  101  Mine          2    2    0    0    0    0
  END GEN-INFO

  PWAT-PARM1
  *** <PLS >          Flags
  *** x - x  CSNO  RTOP  UZFG  VCS  VUZ  VNN  VIFW  VIRC  VLE  IFFC  HWT  IRRG
  101      0    1    1    0    0    0    0    0    0    1    0    0
  END PWAT-PARM1

  PWAT-PARM2
  *** < PLS>  FOREST          LZSN  INFILT  LSUR  SLSUR  KVARV  AGWRC
  *** x - x          (mm)  (mm/hr)  (m)  (1/mm)  (1/day)
  101      0.10  12.36  1.80  15.24  0.15  0.  0.90
  END PWAT-PARM2

  PWAT-PARM3
  *** < PLS>  PETMAX  PETMIN  INFEXP  INFILD  DEEPFR  BASETP  AGWETP
  *** x - x  (deg C)  (deg C)
  101      3.89  0.56  2.  2.  0.  0.02  0.01
  END PWAT-PARM3

  PWAT-PARM4
  *** <PLS >  CEPSC  UZSN  NSUR  INTFW  IRC  LZETP
  *** x - x  (mm)  (mm)  (1/day)
  101      0.00  1.24  0.15  2.00  0.5  0.20
  END PWAT-PARM4

```

```

PWAT-STATE1
*** < PLS> PWATER state variables (mm)
*** x - x      CEPS      SURS      UZS      IFWS      LZS      AGWS      GWVS
      101      0.      12.94      1.27      0.279      12.95      25.4      0.000
      END PWAT-STATE1

END PERLND

PLTGEN

PLOTINFO
Plot-opn ***
# - # FILE NPT NMN LABEL PYR PIVL TYPE ***
1 30 0 16 0
END PLOTINFO

SCALING
Plot-opn ***
# - # YMIN YMAX IVLIN THRESH ***
1 500. 48.
END SCALING

END PLTGEN

EXT SOURCES
<-Volume-> <Member> SsysSgap<--Mult-->Tran <-Target vols> <-Grp> <-Member-> ***
<Name> x <Name> x tem strg<-factor->strg <Name> x x <Name> x x ***
WDM 25 PREC METRZERO SAME PERLND 101 EXTNL PREC
WDM 125 PEVT METRZERO SAME PERLND 101 EXTNL PETINP
WDM 25 PREC METRZERO SAME PLTGEN 1 INPUT MEAN 1
WDM 125 PEVT METRZERO SAME PLTGEN 1 INPUT MEAN 2
END EXT SOURCES

SCHEMATIC
<-Volume-> <--Area--> <-Volume-> <ML#> *** <sb>
<Name> x <-factor-> <Name> x *** x x
PERLND 101 1 PLTGEN 1 91
END SCHEMATIC

MASS-LINK

MASS-LINK 91
<-Volume-> <-Grp> <-Member-><--Mult--> <-Target vols> <-Grp> <-Member-> ***
<Name> <Name> x x<-factor-> <Name> <Name> x x ***
PERLND PWATER SURO 1 PLTGEN INPUT MEAN 3
PERLND PWATER IFWO 1 PLTGEN INPUT MEAN 4
PERLND PWATER UZS 1 PLTGEN INPUT MEAN 5
PERLND PWATER LZS 1 PLTGEN INPUT MEAN 6
PERLND PWATER AGWO 1 PLTGEN INPUT MEAN 7
PERLND PWATER INFIL 1 PLTGEN INPUT MEAN 8
PERLND PWATER PERC 1 PLTGEN INPUT MEAN 9
PERLND PWATER TAET 1 PLTGEN INPUT MEAN 10
PERLND PWATER LZI 1 PLTGEN INPUT MEAN 11
PERLND PWATER UZI 1 PLTGEN INPUT MEAN 12
PERLND PWATER IFWI 1 PLTGEN INPUT MEAN 13
PERLND PWATER SURI 1 PLTGEN INPUT MEAN 14
PERLND PWATER LZET 1 PLTGEN INPUT MEAN 15
PERLND PWATER SURS 1 PLTGEN INPUT MEAN 16
      END MASS-LINK 91

END MASS-LINK

END RUN

```