

Comparison Between Corn Evapotranspiration Rates by the Modified Bowen Ratio and the Ceres-Maize Model

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ABSTRACT

Corn stands out among grains because of its high global importance due to its chemical composition, nutritional value and productive potential. Several factors influence corn crop performance, and climate poses the greatest challenges for crop planning and management. Although tolerant to water deficit, the corn plant presents high sensitivity to water scarcity in specific developmental stages. Therefore, knowing factors related to water loss, namely potential or reference evapotranspiration (ET₀) and Evapotranspiration (ET_c) is crucial. This work aimed to calculate the ET₀ by the methods of Priestley and Taylor (1972) and Penman-Monteith and ET_c of the corn crop, both using the CERES-MAIZE model, and compare them with the results observed by the Modified Bowen Ratio method. The field experiment was conducted in the experimental area of ESALQ / USP and sensors were installed for data collection. For the comparison of results, the CERES-MAIZE model, duly calibrated for the experimental conditions, was used. The results showed that ET₀ was underestimated by the Penman-Monteith method and overestimated by the Priestley and Taylor method through the CERES-MAIZE model throughout the crop cycle. However, at the cycle end, the accumulated values were lower than those measured by the MBR method.

Keywords: Irrigation, climate, water

Introduction

Corn is considered one of the most important crops worldwide (KRESOVIC et al., 2014) due to its chemical composition, nutritional value and productive potential (FANCELLI and DOURADO-NETO, 2000); however, several factors influence its productivity, such as water use. Despite its tolerance water deficit during the vegetative phase, the corn plant shows high sensitivity water scarcity in the stage between pre-flowering and the beginning of grain filling (BERGAMASCHI et al., 2006).

Potential or reference evapotranspiration (ET₀) refers to water losses of a field without water restrictions (PAES et al., 2000) and it can be estimated through different methods, and the method proposed by Penman-Monteith is considered a standard (ALLEN et al., 1998). Crop evapotranspiration (ET_c) is a function of microclimatic conditions of the site of interest, along with the physiological and morphological features of the crop (DOORENBOS and KASSAM 1994;

CARVALHO, 2009), which is important in water consumption estimation. ET_c can be calculated by the Modified Bowen Ratio (MBR) method in which direct measurements of dry and wet bulb temperatures, radiation balance, and soil heat flux are carried out.

Several models are developed to understand the interaction between plants and the environment where they are inserted (OLIVEIRA, 2011). Currently, several efficient models are available, which are tools of great potential in sites of cultivated systems, allowing understanding and estimating crop performance in different locations and conditions (TOJO SOLER, 2004).

Therefore, considering the availability of estimation models for ET₀ and ET_c, this work aimed to calculate the ET₀ by the methods of Priestley and Taylor (1972) and Penman-Monteith and ET_c of the corn crop, both using the CERES-MAIZE model, and compare them with the results observed by the Modified Bowen Ratio method.

Material and Methods

Experimental site

The work was carried out at the Fazenda Areão experimental site, belonging to the Luiz de Queiroz College of Agriculture (ESALQ / USP), in the municipality of Piracicaba, São Paulo State, Brazil (lat 22°52'S and long 47°30'O and altitude 546 m). According to the Köppen classification, the climate of the region is type Cwa, that is, tropical humid, with rainy summers and dry winters. The total rainfall in the driest month (July) is 26 mm and in the wettest month (January), 217 mm. the annual rainfall is about 1270 mm and the driest months are June, July and August. The average temperature in the hottest month (January) is 24.6°C, in the coldest month (July), 17.3°C, with an annual average of 21.5°C.

The soil of the experimental site is classified as Eutrophic Red Argisol and the soil of the experimental site located near the meteorological station is classified as eutrophic Red Nitosol. The soil was prepared in a conventional manner, plowed at a depth of 30 cm, and then trenched for leveling and dewatering. The mineral planting and cover fertilization was carried out according to the soil analysis in order to prevent nutritional stresses. The cover fertilization was carried out in the V5 to V6 stage of the plants, that is, 5-6 fully expanded leaves, with the fertilizer incorporated at 3 cm depth.

The planting spacing was 0.45 m between rows with a population of 66,000 plants per ha. Sowing was performed using about 10% more of the number of seeds due to possible germination failure, pest attack and other germination reduction factors. Hybrid P4285YH was selected, among the most commercialized ones by producers of the region. This hybrid has high productive potential and is well adapted to the edaphoclimatic conditions of the site. Besides an excellent option for silage, grain quality and stalk, hybrid P4285YH has high productivity, leaf stability and sanity as well as high tolerance to storage and breaking.

Modified Bowen Ratio Method

To estimate crop evapotranspiration (ET_c), an automatic modified Bowen ratio system was installed in the center of the experimental site, and was connected to an automatic data acquisition system (datalogger). The system obtained the following meteorological elements: surface radiation balance (R_n); soil heat flux (G) performed at two points (G1 and G2); and vertical gradients of air temperature (T_{ar}) and relative humidity (RH). G1 and G2 sensors were installed 5 cm below the soil surface.

The data were stored daily into the datalogger at 15-min intervals. The energy balance for estimating the heat flux on a surface was defined, with the modified Bowen ratio (β) defined as the ratio

between sensible (H) and latent (LE) heat fluxes. Finally, evapotranspiration was assumed equal to the latent heat flux.

With the measurements of radiation balance (R_n), soil heat flux (G1 and G2), temperature differences (ΔT) and vapor pressure (Δe) between the two levels, energy balance is determined, according to equation 1.

$$R_n - G - H - LE = 0 \quad (1)$$

Where: R_n is the radiation balance (MJ m⁻² d⁻¹), G is the heat flux in the soil (MJ m⁻² d⁻¹), H is the sensible heat flux (MJ m⁻² d⁻¹) and LE is the latent heat flux of evaporation (MJ m⁻² d⁻¹).

The values of the modified Bowen ratio (β) were calculated through equation 2 for each 15-min interval, based on the temperature gradient values (ΔT).

$$\beta = \left(\frac{\Delta T_u}{(1 - W) \Delta T_s} - 1 \right)^{-1} \quad (2)$$

Where: ΔT_s is the temperature difference of the dry bulb, in °C; ΔT_u is the temperature difference of the wet bulb, in °C. E_w is the weighting factor calculated by equations 3 and 4:

$$W = 0.407 + 0.0145 \cdot T_u \quad (0^\circ\text{C} < T < 16^\circ\text{C}) \quad (3)$$

$$W = 0.483 + 0.01 \cdot T_u \quad (16.1^\circ\text{C} < T < 32^\circ\text{C}) \quad (4)$$

The latent heat flux was obtained by equation (5) and the ET_c by equation (6).

$$LE = \frac{R_n - G}{1 + \beta} \quad (5)$$

$$ET_c = \frac{LE}{2.45} \quad (6)$$

The data were analyzed according to methodology developed by Perez et al. (1999) (Table 1), which describes the conditions for the data collected to present physical consistency, avoiding possible estimation errors caused by advection or equipment problems.

Table 1. Conditionals for data evaluation for the modified Bowen ratio method (adapted from PEREZ et al., 1999).

Energy Available	Vapor Pressure Difference	Modified Bowen Ratio	Heat flows
R _n - G > 0	Δe > 0	β > -1	λE > 0 and H ≤ 0 for 1 ≤ β ≤ 0 or H > 0 for β > 0
R _n - G < 0	Δe < 0	β < -1	λE < 0 and H > 0
	Δe > 0	β < -1	λE > 0 and H < 0
	Δe < 0	β > -1	λE < 0 and H ≥ 0 for 1 ≤ β ≤ 0 or H < 0 for β > 0

In periods when inconsistency was observed, it was interpolated based on the previous and posterior values, provided that they did not occur in continuous periods of more than 2 h of incoherent

data. In cases where intervals greater than 2 h did not present consistent data, the entire dataset of the day was discarded.

DSSAT CERES-Maize

The model used was CERES-Maize, grouped in a Decision Support System for Technology Transfer (DSSAT), version 4.6 (HOOGENBOOM et al., 2012), where the experimental data were registered. The experiment was duly registered and the soil profile was constructed according to the chemical and physical analysis of the experimental site. The daily meteorological data of maximum, minimum temperature, radiation balance and precipitation referring to the experiment period for the simulation to be in accordance with actual conditions.

Subsequently, the sensitivity analysis of the model was performed to make the optimistic and pessimistic estimates of a group of variables that influence the model calibration. The simulations were processed considering a scenario of potential production. For this experiment, we evaluated the DKB 333 hybrid on which the DSSAT / CERES-MAIZE model was initially developed, because this hybrid presents similar characteristics to the hybrid used in the experiment.

The DSSAT / CERES-MAIZE model needs to have its coefficients calibrated to produce the corn growth and development (OLIVEIRA, 2015), adjusted to the experimental conditions, considering the average values of the genetic parameters representative of the experimental site. For that purpose, the calibration was performed manually by visual adjustment between the simulated curves and the observation points through the genetic coefficients P1, P2 and P5 that define the crop phenology, while G2 and G3 are related to grain yield, according to Hoogenboom et al. (1994).

The CERES-Maize model can calculate the reference evapotranspiration (ET₀) by two methods, Priestley-Taylor and Penman-Monteith. The Penman-Monteith method is expressed by the following equation (7):

$$ET_0 = \frac{0.408 \cdot s \cdot (R_n - G) + \frac{\gamma \cdot 900 \cdot U_2 \cdot (e_s - e_a)}{T + 275}}{s + \gamma \cdot (1 + 0.34 \cdot U_2)} \quad (7)$$

Where: ET₀ is the evapotranspiration rate of a hypothetical reference crop (mm d-1), R_n is the net surface flux density (MJ m-2 d-1), G is the soil heat flux density (MJ m-2 d-1), γ the psychrometric constant (kPa °C-1), T is the average air temperature (°C), U₂ is the wind speed (ms-1) at 2 m above the soil, e_s is the current vapor pressure of the air (kPa), e_a is the current vapor pressure curve (kPa) and s is the slope of the vapor pressure curve at temperature of the area (kPa °C-1).

Priestley-Taylor Method

The Priestley-Taylor method to calculate daily ET₀ (mm d-1) replaces the aerodynamic term of the Penman-Monteith equation with an empirical dimensionless multiplier (Priestley-Taylor coefficient) thus requiring a smaller data amount. This method is adopted as a standard in the DSSAT software (Equation 8).

$$ET_0 = \frac{1}{\lambda} \cdot \frac{s \cdot (R_n - G)}{s + \gamma} \cdot \alpha \quad (8)$$

Where: (MJ m-2 d-1) is the heat flux of the soil, R_n (MJ m-2 d-1) is the net radiation, G (MJ m-2 d-1) is the latent heat of vaporization, s (kPa °C-1) is the slope of the vapor pressure curve at air temperature, γ (kPa °C-1) is the psychrometric constant, and α is the Priestley-Taylor coefficient. The crop evapotranspiration (ET_c) can be obtained by the relation (Equation 9):

$$ET_c = K_c \cdot ET_0 \quad (9)$$

Where: K_c is the crop coefficient, which varies with the phenological phases as a function of the leaf area index (LAI), as presented in the following equation (Equation 10):

$$K_c = 1.0 + (K_{c \max} - 1.0) \cdot \frac{LAI}{6.0} \quad (10)$$

Evaluation of the results obtained

The MBR was considered an observed evapotranspiration data, where the PM and PT method was the simulated data. The evaluation of its performance was based on the following statistics: correlation (r), root mean squared error (RMSE), BIAS index, concordance index (d) and coefficient of efficiency of the model (COE).

Correlation (r) quantifies the degree of relationship between two variables. The value r varies from -1 to 1, the closer to the upper or lower limits, the greater the relation between estimated and observed values (Equation 11).

$$r = \frac{\sum (E_i - \bar{E})(O_i - \bar{O})}{\sqrt{\sum (E_i - \bar{E})^2 \sum (O_i - \bar{O})^2}} \quad (11)$$

The root mean squared error (RMSE) establishes the mean error of the model, i.e., the lower the value, the better the model performance (MARTINS et al., 2014) (Equation 12).

$$RMSE = \sqrt{\frac{\sum_{i=1}^n (E_i - O_i)^2}{n}} \quad (12)$$

The mean deviation of estimated values based on observed values expresses the model trend, and is quantified by the BIAS index. Thus, the closer to zero, the lower the model trend (Leite, Valdir, Lima et al., 2002; Martins et al., 2014) (Equation 13).

$$\text{BIAS} = \frac{\sum_{i=1}^n E_i - \sum_{i=1}^n O_i}{\sum_{i=1}^n O_i} \quad (13)$$

The concordance index (d) represents much free estimated variables are error-free, and d can range from 0 to 1, that is, the closer to 1, the greater the concordance between observed and estimated data (WILLMOTT, 1981) (Equation 14).

$$d = 1 - \left\{ \frac{[\sum (E_i - O_i)^2]}{[\sum ((E_i - \bar{E}) + [O_i - \bar{O}])^2]} \right\} \quad (14)$$

The coefficient of efficiency of the model (COE) represents how good the adjustment of the simulated values is. COE can vary from $-\infty$ to 1, where the closer to 1, the better the fit in the model (BRAS; CI; SOLO, 2003) (Equation 15).

$$\text{COE} = \frac{\sum_{i=1}^n (O_i - E_i)^2}{(O_i - \bar{O})} \quad (15)$$

Where: E_i is the estimated evapotranspiration values; O_i is the evapotranspiration values observed by MBR; n is the number of observations; \bar{O} is the mean of the observed values; \bar{E} is the average of estimated values.

Results and Discussion

The evaluation of the reference evapotranspiration (ET₀) data estimated from the Pennam-Montheth (PM) and Priestley and Taylor (PT) methods and the data estimated by the Modified Bowen Ratio (MBR) method showed significant differences between the results. Similar observations were found for crop evapotranspiration (ET_c).

According to the results of the statistical analysis (Table 2), the estimated data did not present satisfactory adjustment compared to the measured data, which showed low accuracy. The coefficient of determination (R^2), which occurs in the total variability explained by the model (MONOD et al., 2006), was far from coefficient 1, which indicates a perfect linear correlation. The "d" index of Willmott presented similar behavior to that of the correlation

coefficient; however, the indices for ET_c were slightly higher those for ET₀ in both methods used.

The COE values were low, as they should vary from $-\infty$ to 1, and the closer to 1, the more accurate adjustment (ASCE, 1993). Therefore, the data obtained are not accurate and are considered questionable since they were lower than 0.36 (SILVA et al., 2008).

As observed, for ET₀, estimates using the Pennam-Montheth (PM) method were underestimated (Figure 1) and the results were overestimated (Priestley and Taylor) (Figure 2). For ET_c, the same data behavior was observed for the Pennam-Montheth (PM) method (Figure 3) and for the Priestley and Taylor (PT) method (Figure 4).

The ET₀ estimated by the CERES-MAIZE model using the Pennam-Montheth (PM) and Priestley and Taylor (PT) methods was lower when compared to the ET₀ obtained by the MBR at approximately 90 days after sowing until the end of the corn crop cycle (Figure 5), which may have been due to the small number of meteorological input variables in the model.

With the meteorological input data for the model, it was inserted radiation balance, maximum temperature, minimum temperature and precipitation; However, they were not enough to estimate ET₀ by the Pennam-Montheth (PM) method, since this method still requires de information regarding air humidity and wind speed, which many times cause limitations to its use (PEREIRA et al., 1997). The method of Priestley and Taylor (1972) is more simplistic and approximates the Penman method, using as input variable only maximum and minimum temperature (Figure 6).

Figure 1. Reference Evapotranspiration (ET₀) of the data measured by the Modified Bowen Ratio as a function of the data estimated by ET₀ calculation using the Pennam-Montheth (PM) method.

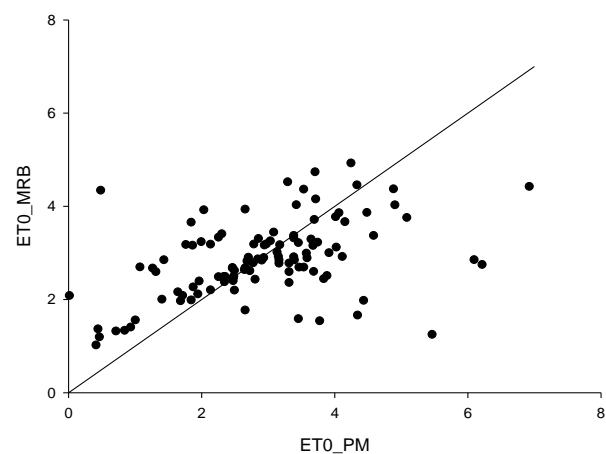


Figure 2. Reference evapotranspiration (ET₀) of the data measured by Modified Bowen Ratio compared to ET₀ calculated according to the data estimated by ET₀ calculation using the Priestley and Taylor (PT) method.

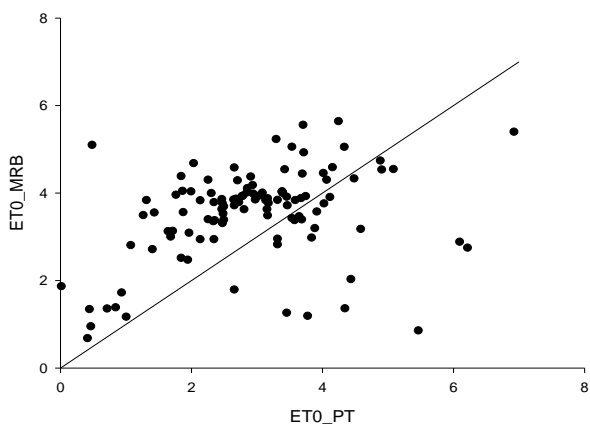


Figure 3. Evapotranspiration of the crop (ET_c) of the data measured by the Modified Bowen Ratio according to the data estimated by ET₀ calculation using the Pennam-Montheth (PM) method.

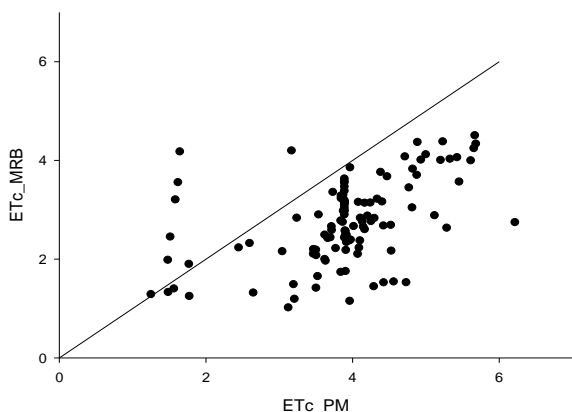
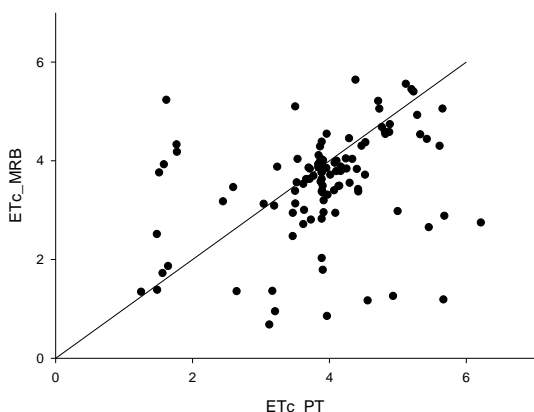


Figure 4. Evapotranspiration of the crop (ET_c) of the data measured by the Modified Bowen Ratio according to the data estimated by ET₀ calculation using the Priestley and Taylor (PT) method.



As observed in the figures below, there was difference between crop evapotranspiration (ET_c) values between both methods used. The ET_c estimated through ET₀ calculated by the Pennam-Montheth model (PM) (Figure 5) was lower when

compared to ET_c measured by the Modified Bowen Ratio (MBR) method (Figure 6). The ET_c estimated through the ET₀ calculated by the Priestley and Taylor (Figure 7) method was superior when compared to ET_c measured by the Modified Bowen Ratio (MBR) method (Figure 8).

Figure 5. Reference Evapotranspiration (ET₀) of the data measured by the Modified Bowen Ratio as a function of days after sowing, by calculating the ET₀ using the Pennam-Montheth (PM) method.

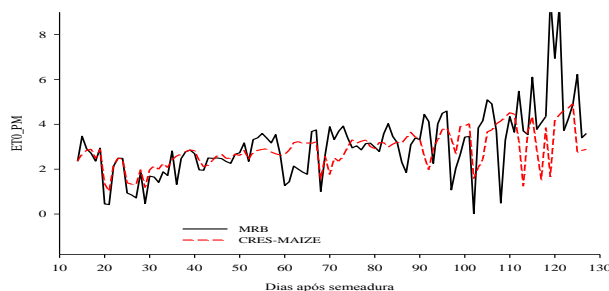
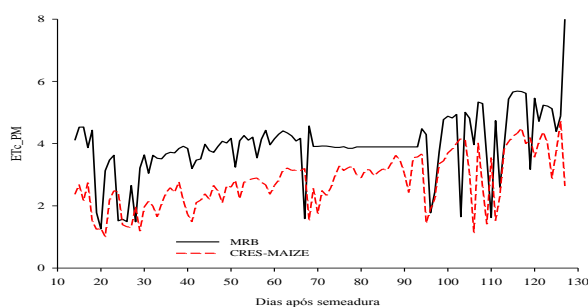


Figure 6. Evapotranspiration of the crop (ET_c) of the data measured by the Modified Bowen Ratio according to the days after sowing, by calculation of ET₀ by the Pennam-Montheth (PM) method.



The cumulative evapotranspiration in the corn crop cycle for the Modified Bowen Ratio (MBR) method was 433.6943 mm d⁻¹, using the Pennam-Montheth (PM) method 308.832 mm d⁻¹ and by the Priestley and Taylor (PT) method, 366.16 mm d⁻¹. However, in spite of ET_c variation throughout the crop cycle, ET₀ estimation using the CERES-MAIZE model were lower than in the accumulated ET_c method (Figure 9).

Figure 9. Evapotranspiration of the crop (ET_c) accumulated during the crop cycle.

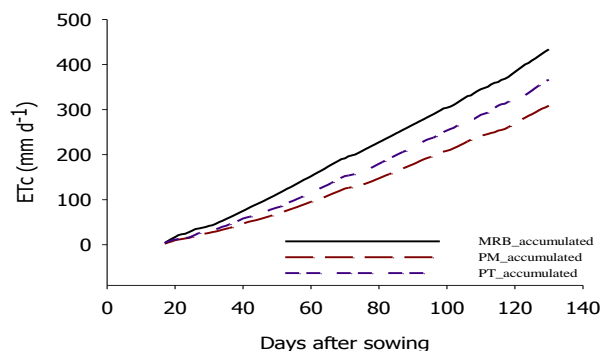


Table 2. Statistical results of reference evapotranspiration (ET₀) and daily evapotranspiration (ET_c) estimated by the Pennam-Montheth (PM) and Priestley and Taylor (PT) methods, using the CERES-MAIZE model in comparison to reference evapotranspiration (ET₀) and crop evapotranspiration (ET_c) estimated by the Modified Bowen Ratio (MBR) method.

		R	R ²	D	RQME	BIAS	COE
PM	ET	0.497		0.531	1.434	-	
	c	5	0.2475	3	6	0.2879	-1.2696
	ET	0.476		0.630	1.400	-	
	o	0	0.2265	0	3	0.0665	0.2068
PT	ET	0.369		0.593	1.357	-	
	c	9	0.1368	4	9	0.1554	-0.8988
	ET	0.314		0.567	1.657	-	
	o	7	0.0991	1	8	0.1502	-0.1117

Conclusions

Potential evapotranspiration (ET₀) was underestimated by the Pennam-Montheth method and overestimated by the Priestley and Taylor method, using the CERES-MAIZE model during the crop cycle.

Accumulated crop evapotranspiration (ET_c) estimated by the Pennam-Montheth and Priestley and Taylor method using the CERES-MAIZE model was lower than that measured by MBR.

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