

Characterization of Hungarian Red Soils

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Red clays in Hungary are the products of soil forming processes occurring during the Tertiary period, not covered later by marine sediments. They were eroded under the climatic conditions of the Quaternary period, so at present they can only be found in sites protected against degradation, or where – due to their thickness and resistance – they could withstand the forces of erosion. Thus, red clays are fossil or relic products of soil formation. Their economic importance is far from negligible; vineyards, forests and arable cultivation can be found on them. Their influence may be detected in larger areas to where they have been washed away, settled and have mixed with other soils.

Different ideas have been formed on the distribution and characteristics of red clays and loamy products (KRETZOI, 1969; BIDLÓ, 1985; JÁNOSSY, 1979; JÁMBOR, 1980; BORSY & SZŐÖR, 1979–1980; SCHWEITZER, 1993). STEFANOVITS (1963, 1967) found that the red clays of Hungary are genetically diverse.

Hungarian red clays differ greatly in their genesis and their physical and chemical characteristics from other soil types in the country, and also from red clays found elsewhere (FEKETE et al., 1997; FEKETE & STEFANOVITS, 2002a,b). Red clays in Hungary are similar to the tropical and sub-tropical ferrallite soils (FEKETE, 1988, 2002) regarding their formation and mineral characteristics. One of our aims is to explore the similarities in processes and characteristics, which would substantially help classification.

Materials and Methods

Samples were collected from nearly two hundred soil profiles from different parts of the North Hungarian Mountains and Transdanubian Mountains regions. 13 samples out of these were selected to represent the different types and sources of red clays (Table 1).

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Table 1
Place of origin of the selected red clay samples

No. of sample and place of origin		Depth in cm	Region	
<i>Red clays in Northern Hungary</i>				
100.	Jósvafő	20–55	North Hungarian Mountains	Aggtelek Karst (limestone)
108.	Mád	40–60	Foreland of North-Hungarian Mountains	Tokaj Mountains rhyolite tuff
222.	Salgótarján	400–450		North-Hungarian Basin
237.	Mátrakeresztes	200–230		Mátra Mountains (volcanic rock 237, 210, 236)
210.	Gyöngyöstarján	20–40		
236.	Muzsla	10–30		
244.	Szurdokpüspöki	350–380		
35.	Hatvan	50–67		Gödöllő Hills
152.	Valkó	260–290		
86.	Gödöllő	30–60		
<i>Red clays in Transdanubia</i>				
119.	Szekszárd	70–80	Transdanubian Hills	Szekszárd Hills (tertiary-quatarnary loam, clay)
120.	Kakasd	60–80		Mecsek Mountains (permian sandstone)
205.	Kővágószőlős	8–15		

Analyses were carried out according to the methodology laid down in BUZÁS (1988, 1993). X-ray diffraction and (derivatographic) thermal analysis were applied to determine the mineral composition of the samples. Micromorphological studies were performed (SZENDREI, 2000). Thin sections were prepared by the method of diluted polyester resin impregnation. The micromorphological features were described according to BULLOCK et al. (1985).

Analytical Results and Assessment

Basic soil analysis

The soil texture varies between medium-heavy silt and heavy clay. The clay content is significantly high, although in red clays its dispersion is also high. The plasticity value (K_A) is in accordance with the hygroscopicity value (hy_1). The highest hygroscopicity values were detected in samples 100-Jósvafő, 210-Gyöngyöstarján, 35-Hatvan.

Table 2
Basic soil analysis data of red clay samples

No.	Place of origin	Depth (cm)	K _A	hy ₁	pH		CaCO ₃ %	Humus %
					KCl	H ₂ O		
<i>Red clays in Northern Hungary</i>								
100.	Jósvafő	20–55	62.00	9.72	4.74	5.92	0.00	0.19
108.	Mád III	40–60	41.00	5.73	6.12	6.79	0.00	0.92
222.	Salgótarján	400–450						
237.	Mátrakeresztes	200–230	68.00	7.50	5.28	6.69	0.00	2.60
210.	Gyöngyöstarján	20–40						
236.	Muzsla	10–30	51.00	3.90	6.86	7.59	0.74	2.34
244.	Szurdokpüspöki	350–380	53.00	2.80	3.77	5.12	0.00	0.32
35.	Hatvan	50–67	96.00	9.42	7.38	8.38	0.45	0.12
152.	Valkó	260–290	47.00	4.70	7.24	8.15	6.32	1.08
86.	Gödöllő	30–60						
<i>Red clays in Transdanubia</i>								
119.	Szekszárd	70–80	57.00	4.18	7.81	8.36	4.14	1.12
120.	Kakasd	60–80	51.00	3.73	7.74	8.36	0.00	0.18
205.	Kövágószőlős	8–15						

K_A = upper limit of plasticity according to Arany; hy₁ = hygroscopicity according to Kuron, modified by Sík

Chemical reaction (pH) is slightly acidic, but some samples have neutral or alkaline chemical reaction, because they contain CaCO₃. The humus content is generally low.

Chemical analysis

The data of total chemical analyses were considered important both for identifying the age of soils as well as for determining the weathering characteristics. The SiO₂/R₂O₃ molecular ratio refers to the nature of weathering. From the total chemical analyses the SiO₂, Al₂O₃ and Fe₂O₃ values are presented in percentage and their ratios for the clay fraction (Table 3).

Based on the molecular ratios the samples can be subdivided into the following genetic groups:

– In the 100-Jósvafő sample the weathering intensity is as strong as in tropical ferrallitic soils. This intensity is shown by the low SiO₂/R₂O₃ ratio; this value is 1.74 in the clay fraction. Similar tropical weathering is indicated in samples 236-Muzsla, 86-Gödöllő, 108-Mád, 119-Szekszárd and 120-Kakasd, where the SiO₂/R₂O₃ ratio of the clay fraction is around 2, or somewhat higher. The Al₂O₃/Fe₂O₃ ratio of these samples ranges between 3.1 and 4.9%.

According to the mineralogical analysis of these samples (Table 4) the kaolinite content is significant, but montmorillonite can be found as well (236-Muzsla).

All above mentioned indicate slightly *ferrallitic weathering*.

Table 3
Chemical analysis data of the selected red clay samples

No.	Place of origin	Depth, cm	in the total sample				in the fine fraction					
			SiO ₂	Al ₂ O ₃	Fe ₂ O ₃	$\frac{\text{SiO}_2}{\text{R}_2\text{O}_3}$	$\frac{\text{Al}_2\text{O}_3}{\text{Fe}_2\text{O}_3}$	SiO ₂	Al ₂ O ₃	Fe ₂ O ₃	$\frac{\text{SiO}_2}{\text{R}_2\text{O}_3}$	$\frac{\text{Al}_2\text{O}_3}{\text{Fe}_2\text{O}_3}$
			%				%					
<i>Red clays in Northern Hungary</i>												
100.	Jósvafő	20-55	50.99	19.54	6.13	3.71	5.03	36.41	30.03	8.78	1.74	5.37
108.	Mád III	40-60	58.21	14.07	4.56	5.84	4.93	40.09	20.79	9.22	2.56	3.54
222.	Salgótarján	400-450						49.84	19.85	7.33	3.46	4.25
237.	Mátrakeresztes	200-230						59.39	28.01	10.20	2.93	4.31
210.	Gyöngyöstarján	20-40	60.31	17.27	7.61	4.57	3.63	50.33	21.41	4.70	3.51	7.14
236.	Muzsla	10-30						48.88	27.24	13.64	2.31	3.13
244.	Szurdokpuszpöki	350-380						50.77	21.41	9.34	3.15	3.60
35.	Hatvan	50-67	59.48	18.80	4.61	4.65	6.34	49.64	21.34	6.98	3.27	4.80
152.	Valkó	260-290	63.32	22.36	4.56	4.86	7.82	53.20	18.49	7.49	3.89	3.87
86.	Gödöllő	30-60	67.77	13.33	3.10	7.52	6.89	39.25	23.97	8.34	2.28	4.51
<i>Red clays in Transdanubia</i>												
119.	Szekszárd	70-80	55.76	30.36	4.97	2.99	9.76	37.35	23.70	9.07	2.15	4.18
120.	Kakasd	60-80	66.96	12.62	4.35	7.39	4.64	38.40	26.64	8.56	2.03	4.98
205.	Kővágószőlős	8-15	69.60	13.62	3.93	7.34	5.56	48.01	25.80	6.36	2.73	6.48

Table 4
Mineral composition of the selected red clay samples (%) in the fine fraction

Minerals	Northern Hungary										Transdanubia			
	Jósvafő (100)	Mád (108)	Salgótarján (222)	Mátrokereztés (237)	Gyöngyös-tarján (210)	Muzsla (236)	Szurdokpüspöki (244)	Hatvan (35)	Valkó (152)	Gödöllő (86)	Szekszárd (119)	Káskasd (120)	Kövágószőlős (205)	
Quartz	37.2	28.1	44.0	30.0	23.0	44.0	52.0	67.1	16.7	30.1	21.6	13.0	56.0	
Calcite	-	1.3	-	-	-	1.0	-	6.3	1.4	5.8	10.6	-	-	
Dolomite	-	-	-	-	-	-	-	-	-	0.6	-	-	3.0	
Feldspars	2.0	13.1	10.0	11.0	7.0	8.0	7.0	-	8.9	4.1	4.7	6.3	7.0	
Kaolinite	30.7	8.1	-	1.0	2.0	-	-	1.0	9.3	9.5	14.8	19.4	-	
Chlorite	-	-	2.0	-	-	3.0	3.0	4.5	1.6	-	10.6	4.0	2.0	
Illite	-	-	6.0	2.0	-	6.0	9.0	-	-	-	-	-	12.0	
Illite + montmorillonite	-	-	5.0	1.0	-	4.0	4.0	-	-	-	-	-	3.0	
Illite + mica	-	-	-	-	-	-	-	5.5	37.9	1.9	-	-	-	
Montmorillonite	-	-	24.0	47.0	50.0	26.0	19.0	-	-	-	-	-	8.0	
Montmorill. + amorphous	23.2	42.8	-	-	-	-	-	-	4.6	40.0	32.0	48.9	-	
Muscovite	-	-	-	-	-	-	-	8.0	-	-	-	-	-	
Gibbsite	-	-	-	-	-	-	-	-	0.8	-	-	-	-	
Hematite	0.5	-	2.0	5.0	12.0	3.0	1.0	-	1.1	-	-	0.8	-	
Goethite	-	-	3.0	-	-	-	-	-	-	-	-	-	5.0	
Humus	2.0	1.4	-	-	1.0	-	-	2.0	1.2	1.6	1.5	1.0	-	
Amorphous	-	-	4.0	5.0	6.0	5.0	5.0	-	7.3	-	-	-	4.0	
H ₂ O ⁻	3.6	4.0	-	-	-	-	-	5.6	3.8	5.4	4.2	-	-	
H ₂ O ⁺	0.8	1.2	-	-	-	-	-	5.6	1.2	1.0	-	-	-	

In conformity with the $\text{Al}_2\text{O}_3/\text{Fe}_2\text{O}_3$ ratio (5.37) the weathering is *allitic* (*bauxitic*) in the 100-Jósvafő sample. Tropical soil features like kaolinite and hematite contents can be found in it.

– In samples 222-Salgótarján, 237-Mátrakeresztes, 244-Szurdokpüspöki, 35-Hatvan and 152-Valkó the weathering is *siallitic*, considering that the $\text{SiO}_2/\text{R}_2\text{O}_3$ ratio falls in the 3–4 range. Their kaolinite content is very small, but montmorillonite (10–47) is significant. The above-mentioned samples are found in the Mátra Mountains, and in the Mátra Foothills, the foreland of the North-Hungarian Mountains region. These samples are red clays, but loess is also involved. In this case the resulted chemical composition shows siallitic weathering.

– In the 205-Kövágószőlős sample the weathering is *siallitic*, because the soil was formed on Permian sandstone.

It is likely that soils (red clays) with higher $\text{SiO}_2/\text{R}_2\text{O}_3$ ratios are resultants of previous tectonic inversion and surface redeposition processes, and thus have a mixed composition of various weathering products (FEKETE & STEFANOVITS, 2002a,b).

Mineralogy

The examination of the mineral composition by thermal analysis and X-ray diffraction is of decisive importance in identifying the age and weathering processes of red clays and also in assessing numerous characteristics of these clays.

On the basis of pedological and mineral analysis the red clays of Hungary can be ranked in various groups:

– Red soils formed on the Permian sandstone: Clay on the surface of Permian red sandstone does not contain kaolinite and hematite, but contains goethite (Sample: 205-Kövágószőlős).

– Red clays on Pannonian surface: In the Transdanubian Hills the red clays have been formed by weathering of the Pannonian surface. Illite, chlorite, montmorillonite and kaolinite can be found (Samples: 119-Szekszárd, 120-Kakasd).

– Red clays formed on rhyolite tuff: The red clays of the Tokaj Mountains, formed on rhyolite tuff with a low amount of kaolinite and 40% montmorillonite in the fine fraction (Sample: 108-Mád).

– Red soils of the Aggtelek Karst, Torna Hills: with 20–30% montmorillonite and 30% kaolinite content (Sample: 100-Jósvafő).

– The red clays in the Mátra Mountains and foothills: These samples have a significant quantity (20–50%) of montmorillonite. It is characteristic of samples 222-Salgótarján, 236-Muzsola and 244-Szurdokpüspöki that they contain little chlorite. There is no chlorite in samples 210-Gyöngyöstarján and 237-Mátrakeresztes. Low hematite contents (1–12%) can be found in them.

– The red soils of the Northern periphery of the Hungarian Plain with significant amount of montmorillonite, and 10% to 20% kaolinite (Samples: 35-Hatvan, 152-Valkó, 85-Gödöllő).

Micromorphology (Table 5)

In the red clays of the Northern periphery of the Hungarian Plain micromorphological features of clay mobilization can be found in the groundmass (speckled, granostriated) in varying degrees as well as along the pores (coatings, hypocoatings and infillings) with different frequency (Samples: 35-Hatvan; 152-Valkó).

In the red clays of the Aggtelek Karst micromorphological features of clay mobilization were noted in the groundmass (speckled, granostriated, monostriated) and along the pores (coatings, hypocoatings and infillings) indicating stresses (Sample: 100-Jósvafő).

In the red clays of the foothills of the Tokaj Mountains: micromorphological features recognized are pronounced in the groundmass (speckled, granostriated, monostriated) and along the pores (coatings and infillings) indicating clay mobilization (Sample: 108-Mád).

Red clays formed by the weathering of the Pannonian surface: Speckled, granostriated and monostriated b-fabrics were only recognized. Clay coatings and infillings were absent probably due to the calcite content (Sample: 119-Szekszárd).

Table 5
Micromorphological features of clay components in the selected red clay samples

Sample		b(birefrangence)- fabric				Pedological features						
						clay/ferruginous clay						
No.	Place of origin	masked b-fabric	mosaic speckled	striated	granostriated	frequency	homogeneous	heterogeneous	coating	infilling	hypocoating	papule
<i>Red clays in Northern Hungary</i>												
100.	Jósvafő		x	x	x	vc			x	x	x	
108.	Mád		x	x	x	c	x		x	x		x
222.	Salgótarján				(x)	c			x	x		x
237.	Mátrakeresztes		x	x	x	vr				x		
210.	Gyöngyöstarján			x	x	vr	x			x		
236.	Muzsola		x	x	x	c	x		x	x	x	x
244.	Szurdokpüspöki		x	x	x	vr			x		x	(x)
35.	Hatvan				(x)	vc	x		x	x		
152.	Valkó		x		x	c	x		x	x	x	x
86.	Gödöllő				(x)	vr			x	x		
<i>Red clays in Transdanubia</i>												
119.	Szekszárd		x	x	x							
120.	Kakasd		x	x	x							
205.	Kővágószőlős		x	x	x	vc		x	x	x		

Remarks: vr: very rare, r: rare, c: common, vc: very common

Red soils formed on Permian sandstone: Well marked micromorphological features indicating mobilization of clay particles in the groundmass (speckled, granostriated, monostriated) and along the pores (coatings and infillings) were observed (Sample: 205-Kővágószőlős).

The speckled and granostriated b-fabric of the groundmass, the occurrence of which is mainly due to swelling and shrinking, was observed in samples 86-Gödöllő, 210-Gyöngyöstarján, 35-Hatvan:Gombos, 100-Jósvafő, 120-Kakasd, 205-Kővágószőlős, 108-Mád, 237-Mátrakeresztes, 236-Muzsla, 222-Salgótarján, 119-Szekszárd, 244-Szurdokpüspöki and 152-Valkó. Swelling was confirmed by the occurrence of stress coatings in the red clays from 100-Jósvafő and 244-Szurdokpüspöki.

Occurrence of clay coatings in samples 108-Mád, 236-Muzsla, 35-Hatvan:Gombos, 205-Kővágószőlős, 152-Valkó, 237-Mátrakeresztes, 222-Salgótarján, 244-Szurdokpüspöki and 86-Gödöllő were interpreted as micromorphological features of illuviation.

Illuviation coatings and infillings were taken notice of most often in samples 35-Hatvan:Gombos and 205-Kővágószőlős.

Summary

Some pedological and micromorphological investigations were carried out on representative red clay samples selected from a large number of profiles. On the basis of conclusions drawn from the analytical results, the red clays can be divided into the following groups:

– The *red clays of the foothills of the Tokaj Mountains* were formed on rhyolite or rhyolite tuff, and are covered by loess in some areas. They are relic soils, older than loess, formed under the warm climate of the Tertiary Period. In addition to quartz they contain feldspars, illite, montmorillonite and a small amount of kaolinite.

– The *red clays of Aggtelek Karst* are Tertiary relic soils formed on Mesozoic limestone. The dominant clay mineral is kaolinite, but they contain a significant amount of smectite as well.

– The *red clays of the Northern periphery of the Hungarian Plain* are situated on clay, silt and sand layers of different origin or between loess depositions. They were formed in the Pliocene and at the turn of the Pliocene and Pleistocene. These soils have a medium clay content, with a large quantity of montmorillonite and a small amount of kaolinite.

– *Red clays formed on Permian sandstones*. These rocks were formed in the Permian period, and were issued from a mixture of sediments under tropical climate, tropical weathering. They are the signs of the oldest soil formation in Hungary. They can be characterized by their kaolinite, illite, montmorillonite and hematite contents.

– The *red clays of the Transdanubian hilly region* were formed by the weathering of the Pannonian surface between the end of the Miocene Period and the lower

Pleistocene. Medium clay content is characteristic of these red clays. They contain kaolinite, montmorillonite, chlorite and a small amount of hematite.

Concerning micromorphological features, speckled and granostriated b-fabrics of the groundmass, mainly due to swelling and shrinking, were observed in some samples. Clay coatings are mainly interpreted as micromorphological features of illuviation.

The investigated red clays are similar to tropical and sub-tropical ferrallitic soils in respect of their formation and mineral characteristics.

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Key words: red clay, relic, fossil, weathering, micromorphological features

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