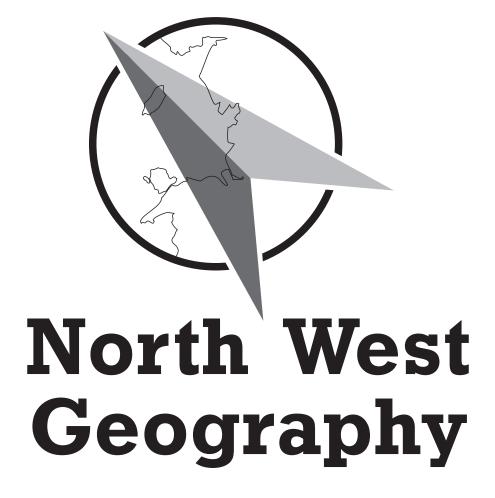
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Sefton Coast's vulnerability to coastal flooding using DEM data

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Abstract

A preliminary analysis of the vulnerability of the Sefton coast to coastal flooding was carried out using a high-resolution DEM and census data from 2011. Results indicate that up to 12,500 people live within areas below the significant dune erosion level. Low-lying areas are clustered in two main locations, the South Formby / Hightown and Southport. High dunes from 15 to 35 m high are common in the central part of the Sefton Coast, including Formby Point and Ainsdale. This highlights the significance of the Sefton dune field system as a potential defence mechanism against coastal erosion and flooding, and the need to consider management schemes that would allow the dunes to adapt to sea level rise and climate change.

Keywords

Dunes, sea-level rise, sustainable management, coastal resilience, low-lying areas, Sefton coast

Introduction

Coastal erosion poses significant threats to coastal populations, particularly those living within a few kilometres of the coast. As much as 70% of the world's sandy coastlines were affected by erosion problems in the second half of the 20^{th} century (Bird, 1985), with recent findings suggesting accelerating erosion rates towards the beginning of the 21^{st} century in many coastal locations (Mars and Houseknecht, 2007; Feagin *et al*, 2005). Despite the negative impacts of coastal erosion to human activities and infrastructure, the number of people choosing to settle at the coast is increasing at rates faster than average national populations (Small and Nicholls, 2003).

The vulnerability of coastal populations has been assessed from a variety of perspectives (McGranagan et al, 2007), with particular attention to their ability to adapt to sea-level rise (SLR) and climate change (Klein and Nichols, 1999; McCarthy, 2001; McGranahan et al, 2007; Nichols et al, 2008) and to the social-environmental factors involved in effective coastal sustainability (Eakin and Luers, 2006; Glavovic, 2006). The protective role of coastal dunes is increasingly being recognized in this context. Coastal dunes are depositional sediment stores that act as protective buffers against coastal flooding (Benavente et al, 2006). They have the ability of migrating landward during periods of SLR while 'holding back the sea' (Davidson-Arnott, 2005) hence providing a low cost, natural mechanism to defend the shoreline against inundation and saltwater intrusion. Recent research including field measurements and modelling demonstrates that under most circumstances sand moves landward with SLR (Aagaard and Sorensen, 2012; Aagaard, 2014) which maintains the overall coastal profile and may even result in shoreline accretion (Houston and Dean, 2014). It is imperative to conduct a detailed assessment of the geomorphic factors that control shoreline evolution along different coastlines and to determine management strategies that would permit adaptation to a slow migration of the dune field inland to accommodate the effects of SLR.

There is no single, universally accepted model to predict shoreline evolution during SLR. Two-dimensional simple methods ignoring onshore sediment movement such as the Bruun Rule (Bruun, 1962) are not adequate and need to be abandoned (Cooper and Pilkey, 2004) in favour of more comprehensive, conceptually valid approaches including complex linkages between sediment transport exchange units in morphodynamic systems (Davidson-Arnott, 2005; Aagaard and Sorensen, 2012; Ranasinghe et al, 2012). The acquisition of airborne scanning laser altimetry (LiDAR) data that can be used to generate digital elevation models (DEM) is used extensively to provide simple flood scenarios under rising sea levels (Brown, 2006; Wang et al, 2002; Webster et al, 2004; Demirkesen et al, 2007; Murdukhayeva et al, 2013). In the absence of well-accepted alternatives to the Bruun Rule there is still a trend towards employing overly-simplified flood models (Cooper and Pilkey, 2004). Despite their significant limitations for appropriate mapping of coastal response GIS and DEMs allow high-quality visualization and

handling of three-dimensional terrains hence overcoming many of the difficulties associated to two-dimensional analysis.

This article provides a preliminary examination of a large LiDAR data set acquired along the Sefton Coast, Merseyside, UK. The coast is relatively well populated and subject to a variety of human uses. Several population centres are fronted by the Sefton dune field, the largest dune field in England, which acts as a protective buffer during severe storms. The first part of the article introduces the generation of a high-quality DEM covering up to 65 km², and uses this to identify near-coast low-lying areas. The second part of the article uses simple census data to provide an indication of how many people live within lowlying sectors. In agreement with previous authors urging to abandon the Bruun Rule and other simplistic approaches to coastal evolution (Cooper and Pilkey, 2004; Davidson-Arnott, 2005), this article does not use low-lying areas to map the spatial extent of potential flood scenarios (i.e., we cannot predict which areas will be inundated based on a DEM). The article does suggest, however, that low-lying areas are likely



Figure 1. Location of study site. Aerial mosaic 2010, courtesy of Sefton Council.

to be more vulnerable, and that there is a potential risk to the people living in them. The aim is to present a relatively quick, preliminary spatial analysis of potentially vulnerable hot-spots in Sefton and hence to demonstrate the increasing significance of adequate coastal dune management along this section of the coast.

Study Site

The Sefton coast extends for 36 km and is influenced by processes in the Irish Sea, the Ribble and Mersey estuaries to north and the south respectively, and the West Lancashire coastal plain (Figure 1).

The coast is macro-tidal with a mean spring tidal range of 8 m. It includes a diversity of coastal environments including salt marshes, beaches, tidal flats, and coastal dunes. Coastal dunes, in particular, extend 16 km alongshore and 4 km inland, covering a total area of 2100 ha and constituting the largest coastal dune field in England, and one of the largest in the UK (Esteves et al, 2012). The Sefton Coast has been internationally designated as EUSAC (Special Area of Conservation), RAMSAR, SSSI, and has received multiple national and local recognitions (Worsley et al, 2012). The dunes have large conservational and economic value and they are subject of multiple activities and intense human pressure (Pye and Blott, 2010). These include agricultural land, important recreational sites for tourists and local residents, and relatively large conservation areas managed by a number of institutions such as Natural England and the National Trust. Beach and dune sediments consist of uniform, well-sorted, fine to medium quartz sands and predominant winds are from the SW (Plater and Grenville, 2010; Esteves et al, 2012).

Methods

DEM

The use of Digital Elevation Models (DEMs) in Geographic Information Systems (GIS) is widely used. Applications include hydrologic modelling (Vieux, 2001), analysis of erosion-deposition processes (Mitasova *et al*, 1996), habitat mapping (Aspinall and Veitch, 1993) and landslide hazards (Olmacher and Davis, 2003) among others. DEMs have also been used to map coastal lowlands (Gambolati *et al*, 2002; Brown, 2006; Wang *et al*, 2002). The procedure often involves in using supervised classifications to identify elevations below a given threshold and the digitalization of low-lying areas into vectors depicting the limits of potential floodplains (Webster *et al*, 2004).

A total of 79 1x1m spot resolution LiDAR raster files were included in the analysis presented here. The

data were collected in 2013. Each raster covered an area of approximately 1.26 km² and was referenced into British National Grid. A new raster mosaic was created from the individual raster files. Focal statistics using rectangle as the neighbourhood and mean as the statistic were employed to correct for NoData values in the original LiDAR files. The UK National Oceanography Centre (NERC) works with tidal levels relative to chart datum (CD). The Sefton DEM was corrected from D_OSGB_1936/ODN datum to CD using the values suggested by the National Tidal and Sea Level Facility (NTSLF) at Formby point. (<u>http://www.ntslf.</u> org/tides/datum). This allows comparison of DEM elevation data with tidal ranges at the site. The resulting DEM was extruded in ArcScene (Figure 2) and basic statistics were calculated (see insert in Figure 2). The average elevation of the dune toe was calculated by analysing the beach-dune boundary (BDB) from Ainsdale to Hightown. Appropriate spatial identification of the dune toe is extremely difficult as this area is characterized by high temporal/spatial variability and the sporadic presence of embryo dunes. As a proxy, the seaward extent of vegetation was used to digitize the beach-dune boundary using common editing procedures. The 3D Analyst toolbox was used to locate outliers and remove them from the beachdune boundary shapefile prior to calculating statistics. The average elevation of the beach-dune boundary along the Sefton coast was 10.32 ± 0.4 m CD. This is approximately 0.30 m above the average dune toe height reported by Esteves *et al* (2012), who observed a high degree of seasonal variability associated with it.

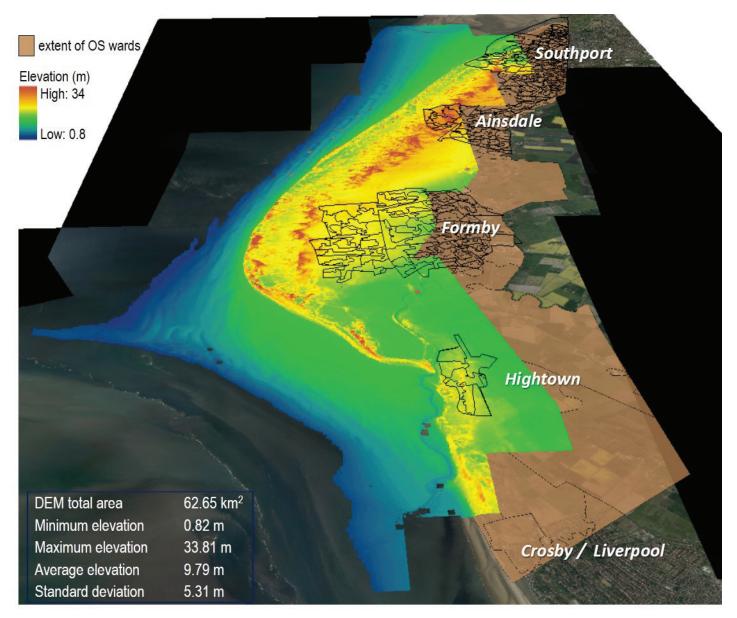


Figure 2. Extent of Sefton DEM and the OS selected wards. Only the seaward sections of Southport, Ainsdale, and Formby fell within the area covered by the LiDAR flight. Hightown was completely within the DEM but none of Crosby's wards could be analysed. Vertical exaggeration x5. Inset includes basic descriptive statistics for the DEM.

Esteves *et al* (2012) classified the type of wave events that result in significant dune erosion (SDE) along the Sefton Coast. Erosion was mainly experienced during periods of peak water levels of 10.2 m (CD) and 2.6 m waves (combined elevation = 12.8 m), rather than during extreme surges and wave heights. The SDE and BDB elevations, alongside tidal elevations provided by the NTSLF for Formby (Table 1), were used to generate binary maps showing the location of low-lying areas in the Sefton DEM.

Table 1. Elevations used to identify low-lying areas in the Sefton coast (based on CD). The significant dune erosion (SDE) value has been adopted from Esteves et al (2012). The beach-dune boundary (BDB) value was estimated in this study. The HAT, MHWS, and MHW values were obtained from the NTSLF at Formby. Water heights are in m.

Elevation	Definition	Height (m)
SDE	Combined effect of 10.2 (water level) + 2.6 wave heights	12.80
BDB	Average height of the Beach-Dune Boundary	10.32
HAT	Highest Astronomical Tide	10.13
MHWS	WS Mean High Water Spring	
MHW	Mean High Water	8.15

Vulnerability Assessment

Wards boundary data were downloaded from the Office for National Statistics website (<u>http://www.ons.gov.uk/</u>) for 2011. This contains Ordnance Survey boundary line data updated annually for each ward in the UK. Cesus data for the selected wards were downloaded from InFuse 2011, the UK Data Service Census Support (<u>http://infuse2011.mimas.ac.uk/</u>). Analyses were conducted at the Output Areas and Small Areas Local Authorities level, the lowest geographical level at which census estimates are provided. Both files were joined and clipped in ArcMap to get a subset of the study area, with particular interest in the main population centres (Southport, Formby, and Hightown). Additionally, the census shapefile was clipped to match the spatial extent of the LiDAR data. This allowed using intersecting tools to provide an estimate of the number of people and urban zones within low-lying areas identified in the DEM.

Results

Figure 3 includes binary maps showing the location of lowlying areas based on elevations in Table 1.

The binary maps show that there are certain areas to the south of Formby and around Hightown where the terrain is below the mean high water (MHW; 8.15 m CD). Most of the beach to the north and south of Formby is above this level but below the mean high water during spring tides (MHWS; 9 m CD). Large areas around Hightown and along the south coast of Southport are at an elevation lower than the highest astronomical tide (HAT; 10.13 m CD) or the beach-dune boundary (BDB; 10.32 m CD). There are multiple sections of the coast below the significant dune erosion level (SDE; 12.8 m CD) including most of the areas south of Formby, a significant amount of frontal dunes in Ainsdale, and the seaward side of Southport. Table 2 includes

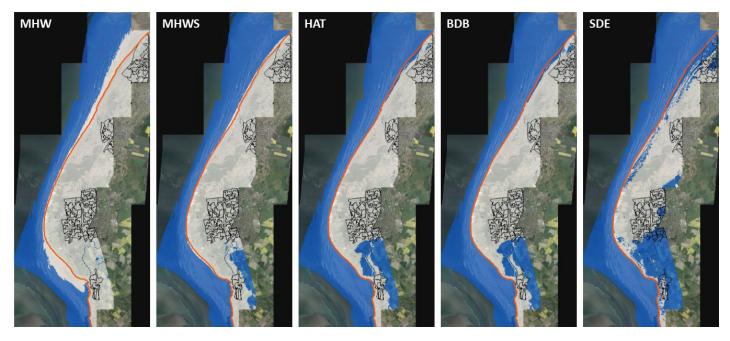


Figure 3. Raster grids showing the location of low-lying areas in Sefton (in blue), defined as areas below the mean high water (MHW), mean high water spring (MHWS), highest astronomical tide (HAT), beach-dune boundary (BDB) and the significant dune erosion (SDE) level identified by Esteves et al. (2011). The spatial location of the BDB used in this study is shown in red.

the areas below and above each elevation and the percent of decrease relative to the MHW map.

Table 2. Areas (km^2) below and above elevations shown in Figure 3. Total area covered by DEM = 62.65 km².

Elevation	Area below	Area above	Decrease (%) of land over particular elevations (relative to MHW)
MHW	24.94	37.70	-
MHWS	31.07	31.57	9.78
HAT	34.73	27.92	15.62
BDB	35.12	27.53	16.24
SDE	41.15	21.50	25.87

Results from overlaying topography and census data indicate that a relatively large number of people live within the low-lying areas (Table 3). The total amount of urban area below the MHW is only 0.65×10^{-3} km² including approximately 929 people (2.87% of the total population analysed in this study) in 2011. The total urban area and number of residents within low-lying areas based on MHWS, HAT, and BDB levels gradually increase up to 0.19 km² and 2679 people (8.27% of residents) for the latter case. There is a dramatic increase in the area and amount of people living below SDE elevations, with almost 2.20 km² of urban zones and 12,500 people (38.56% of the population).

Table 3. Urban areas classified as low-lying areas and number of people living within them. Percentages have been calculated with respect to the total amount of population considered in this study (32,393 residents) as given by the OS census data clipped to the spatial extent of the Sefton DEM (see Figure 2).

Elevation	Urban area (km ²)	Number of people	% population below particular elevations
MHW	0.65 x 10 ⁻³	929	2.87
MHWS	7.72 x 10 ⁻³	1,248	3.85
HAT	95.31 x 10 ⁻³	1,923	5.94
BDB	0.19	2,679	8.27
SDE	2.20	12,492	38.56

Discussion

Potential risks coastal populations and sea level rise

It is worth stressing that maps shown in Figure 3 do not depict modelled flood scenarios. Water levels have been used here to identify relative low-lying areas, not necessarily areas that will be flooded during high water periods. Flooding and coastal morphological changes resulting from increases in sea level strongly depend on time and location, and are regulated by complex mechanisms and non-linear relations between many factors not considered here (e.g., presence of vegetation, sediment transport paths, wave and wind action, duration and timing of storm events, etc.) (Esteves *et al*, 2012). The existence of low-lying areas is a strong indication, however, of the potential vulnerability of sections of the Sefton Coast to future SLR and climate change. The latest 5th Assessment Report (AR5) of the Intergovernmental Panel on Climate Change (IPCC) suggests a number of sea level rise projections depending on different gas emissions scenarios. Predictions include increases in average sea level of approximately 0.2 m (best case scenario) and 0.25 m (worst case scenario) by 2050, and 0.44 m (best case scenario) or 0.73 m (worst case scenario) by 2100. Because of the levels of uncertainty considered by the IPCC sea level could in fact rise even more by up to 1m towards 2100, posing real challenges for the Sefton Coast.

The importance of dune mobility and implications for management

Figure 4 shows the location of large dune areas along the Sefton Coast. While Ainsdale village partially sits over one of the areas including some of the highest dunes in the system, Formby town is almost entirely fronted by very large dunes seaward. This provides in principle a very significant buffer against coastal erosion and flooding. However, the effectiveness of coastal dunes as a defence mechanism strongly depends on a number of variables, including their ability to move and respond to environmental changes (Pye *et al*, 2014).

As in many other locations worldwide (Jackson and Cooper, 2011; Miot da Silva, 2013; Provoost et al, 2009) the Sefton dunes have undergone progressive over-growth of vegetation and re-sealing (Smith, 2012). The loss of bare sand has significant consequences for associated habitats and ecosystems diversity (Smith, 2012). Hence dune regeneration programs are now seeing vegetation stripping and dune remobilization as necessary elements for sustainable dune management (Terlouw and Slings, 2005; Walker et al, 2013; Houston and Dargie, 2010; Pye et al, 2014). The reduction of dune mobility as a result of the over-growth of vegetation also has consequences for dune morphodynamics (Saye and Pye, 2007). However, the potential effects of sea level rise at Sefton should be assessed in the context of the littoral sediment budgets and coastal evolution, particularly in the extent and location of sediment sinks.

Dune erosion may be accentuated in sections of the Sefton dunes prone to wave undercutting during high water levels (Pye and Blott, 2008; Halcrow, 2009; Brown *et al*, 2010; Esteves *et al*, 2012). Complications may arise in areas that were subject to sand mining before and after the Second



Figure 4. Location of Sefton highest dunes. Elevation is in metres.

War World and/or where sea defenses exist (Smith, 2012). Dune erosion during a storm may not be a problem if there is no dune breaching. In fact, cliffing of the foredune may increase the opportunity for landward transfer of eroded material over the dune crest and hence inland migration of the dune to keep pace with rising sea level (Davidson-Arnott, 2005). Recent studies suggest that under most conditions coastal profiles migrate landward with increasing sea-level (Aagaard and Sorensen, 2012; Aagaard, 2014; Houston and Dean, 2014). Higher water levels at Sefton will likely lead to dune erosion but this may facilitate sand transport inland and thus landward migration of the protective dune system. Appropriate analysis of the Sefton coast dynamics at a variety of scales is necessary to investigate sustainable long-term management including strategies that call for managed retreat with preservation of dunes as a response to sea level rise and climate change.

Conclusions

This study presents a DEM that permits some preliminary assessment of the problem of coastal vulnerability at the Sefton Coast. Further studies are required to assess in detail the geomorphic factors that control shoreline evolution and determination of a management strategy that would permit adaptation to a slow migration of the dune field inland to accommodate the effects of the geomorphic evolution and SLR, except in those areas where the littoral budget is sufficiently positive to allow stability of the shoreline or even progradation despite higher water levels. Current understanding of shoreline evolution suggests that there is no need for sand from the dunes to be used to build up the nearshore under rising sea levels (as the Bruun model calls for) (Aagaard and Sorensen, 2012; Houston and Dean, 2014). Contrary to this the volume of dunes can be maintained as they move inland and may be even increased through the addition of some eroded material (Davidson-Arnott, 2005). Hence both dune mobility and space for the coastal dune field to adjust to new conditions are key aspects contributing to the resilience of the coastal dune field, and its ability to act as a sustainable coastal defence.

The Sefton DEM shows that while some of the largest dunes are now completely stabilized below urban centres (e.g., Ainsdale) with sections of the coast prone to coastal squeeze, considerable areas of the dune field have the potential to act as significant buffers against coastal flooding (e.g., Formby point). The existence of low-lying areas clustered around Hightown/South Formby and Southport needs to be further investigated. In particular there is a need to understand what geomorphic factors control the actual vulnerability of these areas in relation to the dune field. Analysis should be conducted to distinguish between areas that are vulnerable to flooding under existing conditions from those that are potentially vulnerable in the absence of high dunes. This could provide the rational to guide dune management and long-term actions designed to enhance dune growth and long-term sustainability of the Sefton dune field.

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