Investigation of Extractable Organic Compounds in Deep-sea Hydrothermal Vent Fluids along the Mid-Atlantic Ridge Thomas M. McCollom¹, Jeffrey S. Seewald², & Christopher R. German² ¹Laboratory for Atmospheric and Space Physics, Campus Box 600, University of Colorado, Boulder CO 80309, email: mccollom@lasp.colorado.edu ²Woods Hole Oceanographic Institution, Woods Hole, MA 02543

10

11 **Abstract** – The possibility that deep-sea hydrothermal vents may contain organic compounds 12 produced by abiotic synthesis or by microbial communities living deep beneath the surface has 13 led to numerous studies of the organic composition of vent fluids. Most of these studies have 14 focused on methane and other light hydrocarbons, while the possible occurrence of more 15 complex organic compounds in the fluids has remained largely unstudied. To address this issue, the presence of higher molecular weight organic compounds in deep-sea hydrothermal fluids was 16 17 assessed at three sites along the Mid-Atlantic Ridge that span a range of temperatures (51 to 18 >360 °C), fluid compositions, and host-rock lithologies (mafic to ultramafic). Sample were 19 obtained at several sites within the Lucky Strike, Rainbow, and Lost City hydrothermal fields. 20 Three methods were employed to extract organic compounds for analysis, including liquid:liquid 21 extraction, cold trapping on the walls of a coil of titanium tubing, and pumping fluids through 22 cartridges filled with solid phase extraction (SPE) sorbents. The only samples to consistently 23 yield high amounts of extractable organic compounds were the warm (51-91 °C), highly alkaline 24 fluids from Lost City, which contained elevated concentrations of C_8 , C_{10} , and C_{12} *n*-alkanoic 25 acids and, in some cases, trithiolane, hexadecanol, squalene, and cholesterol. Collectively, the 26 C_8 - C_{12} acids can account for about 15% of the total dissolved organic carbon in the Lost City 27 fluids. The even-carbon-number predominance of the alkanoic acids indicates a biological 28 origin, but it is unclear whether these compounds are derived from microbial activity occurring 29 within the hydrothermal chimney proximal to the site of fluid discharge or are transported from 30 deeper within the system. Hydrothermal fluids from the Lucky Strike and Rainbow fields were 31 characterized by an overall scarcity of extractable dissolved organic compounds. Trace amounts 32 of aromatic hydrocarbons including phenanthrenes and benzothiophene were the only 33 compounds that could be identified as indigenous components of these fluids. Although 34 hydrocarbons and fatty acids were observed in some samples, those compounds were likely 35 derived from particulate matter or biomass entrained during fluid collection. In addition, extracts 36 of some fluid samples from the Rainbow field were found to contain an unresolved complex 37 mixture (UCM) of organic compounds. This UCM shared some characteristics with organic 38 matter extracted from bottom seawater, suggesting that the organic matter observed in these 39 samples might represent seawater-derived compounds that had persisted, albeit with partial 40 alteration, during circulation through the hydrothermal system. While there is considerable 41 evidence that Rainbow and Lost City vent fluids contain methane and other light hydrocarbons 42 produced through abiotic reduction of inorganic carbon, we found no evidence for more complex organic compounds with an abiotic origin in the same fluids. 43

44 **1. INTRODUCTION**

45 Ever since deep-sea hydrothermal systems first began to be explored in the late 1970s, the 46 occurrence of organic compounds in the hydrothermal fluids has been the subject of considerable 47 scientific study. One of the primary motivations behind these studies has been to determine the 48 source of the organic compounds and, in particular, to investigate whether some compounds 49 might be generated by abiotic organic synthesis within the hydrothermal system (e.g., Welhan and Craig, 1983; Lilley et al., 1983, 1993; Berndt et al., 1996; Charlou et al., 2000, 2002, 2010; 50 51 Holm and Charlou, 2001; McCollom and Seewald, 2001, 2007; Foustoukos and Seyfried, 2004; 52 Proskurowski et al., 2008; Bradley and Summons, 2009; Lang et al., 2010, 2012). In addition, 53 some organic compounds in deep-sea hydrothermal fluids may be produced by biological 54 processes in subsurface environments, and can potentially provide insights into the activity of 55 deep-seated microbial communities (Deming and Baross, 1993; Bradley and Summons, 2009; 56 Reeves et al., 2014).

57 Field-based studies of organic compounds in deep-sea hydrothermal fluids have 58 predominantly focused on the abundance and isotopic composition of methane and light hydrocarbons (e.g., Welhan and Craig, 1983; Lilley et al., 1983, 1993; Charlou et al., 1996, 59 2000, 2002: Proskurowski et al., 2008; Foustoukos et al., 2009). However, several studies have 60 61 analyzed deep-sea hydrothermal fluids for the presence of other dissolved compounds including small carboxylic and amino acids (Haberstroh and Karl, 1989; Martens, 1990; Horiuchi et al., 62 2004; Lang et al., 2010, 2013; Klevenz et al., 2010; Fuchida et al., 2014). Relatively few studies 63 have examined whether higher molecular weight hydrocarbons and other larger extractable 64 organic compounds are present in deep-sea hydrothermal fluids (Brault et al., 1988; Holm and 65 Charlou, 2001; Konn et al., 2009, 2012). 66

Sampling deep-sea hydrothermal fluids for higher hydrocarbons and other organic 67 68 compounds presents significant challenges. Based on the abundance of volatile light hydrocarbons, higher molecular weight compounds might be expected to be present at 69 70 nanomolar or lower concentrations, requiring them to be concentrated prior to analysis for most 71 commonly used analytical methods. In addition, hydrocarbons and many other types of organic 72 compounds must be extracted from the aqueous phase prior to analysis for conventional 73 methods. While hydrocarbons and other low-polarity organic compounds have relatively high 74 aqueous solubilities at the elevated temperatures and pressures that occur within many deep-sea 75 hydrothermal systems, cooling of discharged fluids to ambient temperatures of the deep sea 76 (usually ~2 °C) may cause these compounds to exsolve and precipitate on the surfaces of the 77 sampling device. If only the fluid phase is analyzed, exsolved compounds may not be detected. Furthermore, organic compounds from other sources, including sweater and hydrothermal vent 78 79 mineral deposits (e.g., Simoneit et al., 2004), can be inadvertently entrained into the sample 80 during fluid collection. In addition, hydrothermal vent chimneys are often inhabited by 81 biological communities, enhancing the potential for contamination from biomass if solids are 82 incorporated during sample collection. Oils, hydraulic fluid, and other organic compounds 83 leaking from the vehicle used to collect samples represent other potential sources of 84 contamination.

This study was undertaken to investigate the occurrence of extractable organic compounds in hydrothermal systems along the Mid-Atlantic Ridge (MAR). Hydrothermal fluids were sampled at three sites along the MAR representing a range of conditions (Table 1). Hot (up to 366 °C), acidic fluids were sampled at systems hosted in ultramafic rocks at Rainbow and in basaltic rocks at Lucky Strike (Von Damm et al., 1998; Charlou et al. 2000, 2002). Warm (~90 °C),

90 strongly alkaline fluids circulated through serpentinized ultramafic rocks were sampled at Lost 91 City (Table 1) (Kelley et al., 2005). The Rainbow and Lost City were targeted for study because 92 previous investigations have indicated that the methane and light hydrocarbons in the 93 hydrothermal fluids at these sites may have an abiotic origin (Charlou et al., 2002, 2010; 94 Proskurowski et al., 2008), and it has been suggested that other hydrocarbons with an abiotic 95 origin might also be present (Holm and Charlou, 2001; Konn et al., 2012). None of the systems 96 included in this study show any evidence for buried sediments that might contribute organic 97 compounds to circulating hydrothermal fluids.

98 Concurrent with our study, Konn et al. (2009, 2012) also assessed the presence of extractable 99 organic compounds in hydrothermal fluids from the Rainbow and Lost City sites, but used 100 somewhat different methods for extraction and analysis. They reported detection of a number of 101 extractable organic compounds in the fluids from both sites, including aliphatic, cyclic, and 102 aromatic hydrocarbons as well as fatty acids. Previous work by Holm and Charlou (2001) had 103 identified trace amounts of a homologous series of C₁₆-C₂₉ n-alkanes in hydrothermal fluids at 104 Rainbow. Other studies have identified formate, acetate, and amino acids in the fluids at Lost 105 City (Lang et al., 2010, 2013). Reeves et al. (2014) report concentrations of methanethiol in hydrothermal fluids at many of the same sites included in this study, from samples collected on 106 107 the same expedition as those reported here.

108 **2. METHODS**

109 **2.1. Fluid sampling**

110 Hydrothermal fluid samples were acquired in July 2008 using the Jason remotely operated vehicle (ROV) as part of the KNOX18RR expedition to the Mid-Atlantic Ridge on board the 111 112 R/V Roger Revelle (Fig. 1). Hydrothermal fluids were sampled at several vent sites across the 113 Lucky Strike, Rainbow, and Lost City hydrothermal fields, as summarized in Table 1. A single 114 vent fluid sample was also obtained at the TAG hydrothermal site before a shipboard equipment 115 failure prematurely terminated the cruise, but because no organic compounds were observed in 116 the sample it is not discussed further here. At Lucky Strike and Rainbow, fluids were collected 117 from a number of discrete vent locations distributed across the hydrothermal fields. Most 118 samples were obtained at natural orifices in chimney structures where there was vigorous 119 discharge of high-temperature (>250 °C) hydrothermal fluids (e.g., Fig. 1a). For the Ecurie site 120 at Rainbow, however, fluids were sampled from a cavity excavated into the side of a chimney structure that had moderately hot fluids (~60 °C) slowly venting from its apex (Fig. 1b). 121 122 Theoretical studies have suggested that mixing zones within deep-sea hydrothermal systems 123 might be particularly favorable environments for abiotic organic synthesis (Shock and Schulte, 124 1998), and this site was chosen because mixing of high temperature fluids with seawater is a 125 likely reason for the low temperature of the vent fluids. Since the fluids venting at the apex of the Ecurie structure were too diffuse to sample, a cavity was excavated in the side of the chimney 126 127 to gain access to more vigorous fluid flow (Fig. 1b). At Lost City, fluids were obtained at the 128 Beehive site on the side of a large structure (~60 m high) named Poseidon (Fig. 1c,d) and at an 129 unnamed orifice at the top of the same structure. For comparison with the vent fluids, samples of 130 deep seawater were also collected adjacent to the Lucky Strike, Lost City, and TAG sites using 131 the same methods employed to collect hydrothermal fluids.

Because it was not clear at the outset what approach might be most effective for extraction of dissolved organic compounds from hydrothermal fluids, three different sampling strategies were employed. First, organic compounds were extracted from hydrothermal fluids acquired using

135 isobaric gas-tight (IGT) syringe-type samplers (Seewald et al., 2002). A key feature of the IGT 136 samplers is that the fluid intake rate can be regulated to minimize the amount of seawater 137 entrained during sample collection. In an effort to minimize the contribution of organic 138 compounds from background sources, the original design of the samplers was modified by 139 replacement of the Viton O-rings on the sample chamber piston with Teflon O-rings to eliminate 140 the need for a lubricant. Nevertheless, it was still necessary to use a Viton O-ring and a lubricant 141 (Fluorolube) on the pressure seal for the sample chamber. Prior to each deployment, all wettable 142 surfaces in the samplers were washed with successive rinses of methanol, dichloromethane 143 (DCM), and hexane, and then dried in air. Despite these precautions, traces of fluorolube were 144 found in some samples, presumably derived from the lubricant used on the pressure seal or from 145 residual compounds in the sampler snorkels. The dead volume of the samplers (including the 146 snorkel; ~ 5 ml) was filled with deionized water (Fisher Scientific Optima, hereinafter referred to 147 as Fisher H₂O) prior to deployment rather than the usual practice of filling this volume with 148 bottom seawater. A Teflon cap was placed on the inlet of the sampler snorkel to limit seawater 149 infiltration during transport to the seafloor.

150 In the second approach, hydrothermal fluids on the seafloor were pumped through a 4 m long 151 coil of narrow titanium tubing $(0.125'' \text{ O.D.} \times 0.064'' \text{ I.D.})$ to extract organic compounds (this apparatus came to be known as the Peristaltic Organic Pump sampler or "POP gun") (Fig. 2). 152 153 The concept behind this device was that, because the Ti-coil is bathed in cold (~2 °C) seawater, 154 cooling of the hydrothermal fluids as they passed through the tubing might induce precipitation 155 of higher molecular weight organic compounds such as long-chain hydrocarbons and fatty acids 156 on the tubing walls, owing to the decreasing solubility of these compounds with decreasing temperature and the large internal surface area of the narrow coiled tube. Prior to each 157 158 deployment, the Ti tubing was washed with DCM followed by MeOH, and then rinsed and filled 159 with Fisher H₂O.

160 The third approach used to extract organic compounds was to attach a purpose-built cartridge 161 containing a solid-phase extraction (SPE) sorbent to the outlet of the Ti tubing of the POP gun 162 (Fig. 2 inset). The SPE cartridges were intended to extract additional organic compounds that remained dissolved in the fluid after passage through the Ti tubing. The SPE cartridges 163 164 contained several grams of SPE microbeads within a 2.54 cm outside diameter Teflon tube (2.5 165 cm O.D. \times 8 cm long), capped with porous Teflon frits to allow flow of fluids through the cartridge, and sealed with Swagelok fittings. Prior to deployment, the SPE sorbent was rinsed 166 167 with ~15 ml MeOH followed by ~30 ml of Fisher DI water while in the cartridge.

Two types of SPE sorbents targeted at different classes of organic compounds were used 168 169 separately during deployments of the POP gun (Table 1). For most deployments, the 170 "hydrophilic-lipophilic balanced" Oasis HLB sorbent (Waters Corporation) was used. This 171 sorbent had been used previously for extraction of organic compounds from fluids collected at the Rainbow site (Holm and Charlou, 2001), and is designed to extract a broad spectrum of 172 173 acidic, basic, and neutral organic compounds from aqueous fluids. At two sample locations 174 (Rainbow and Lost City) fluids were extracted using the C8 phase from Grace Discovery 175 Sciences. This phase is designed to target smaller non-polar organic compounds in the octane range, and was used in an attempt to isolate the type of alklylated monocyclic compounds that 176 177 had been reported in previous SPE extracts of hydrothermal fluids (Konn et al., 2009, 2012).

When fully assembled on *Jason*, the POP gun sampler consisted of a coiled 4 m length of titanium tubing connected at the outlet end through Teflon tubing to the cartridge containing the SPE sorbent, and then to a peristaltic pump (Fig. 2). At the seafloor, the inlet of the POP gun was 181 inserted into the orifice of a hydrothermal vent, and fluid was then pumped through the device 182 for a period of about an hour in most cases. During sampling, flow of hydrothermal fluid 183 through the POP gun was visually confirmed by the presence of shimmering (i.e., warm) fluid at 184 the outlet of the peristaltic pump. A benchtop test with the fully assembled POP gun indicated that the flow rate was about 21 ml min⁻¹, although this flow rate may not have been achieved by 185 the peristaltic pump during operation at in situ pressures on the seafloor. Assuming the benchtop 186 187 flow rate to be representative, it is estimated that ~ 1.2 L of hydrothermal fluid would have been 188 pumped through the POP gun during each 60 minute sample. Owing to logistical considerations, 189 POP gun and SPE samples could not be obtained at every site where fluids were collected using 190 IGT samplers during the cruise (Table 1).

191 **2.2. Sample processing**

192 The three sampling strategies employed required use of different protocols to extract organic 193 compounds (Fig. 3). Following deployment and return to the surface, all sampling devices were 194 immediately placed in a refrigerator at 4 °C until they were processed (usually within ~2 hours). 195 All solvents used to extract organic compounds were Fisher Scientific Optima grade.

196 For the IGT samplers, the first two aliquots of fluid from each sample were collected in pre-197 cleaned vials and stored at -20°C for later analysis of inorganic chemistry (5 ml; Table 1) and 198 dissolved free amino acids plus other organic compounds (~20 ml). The remaining fluid (~120 199 ml) was then transferred to a pre-cleaned glass bottle, where organic compounds were extracted 200 by adding ~8 ml of dichloromethane (DCM) and then shaking vigorously for 10 minutes. The DCM phase was then removed by pipette into a separate vial for further processing. Because the 201 202 Lost City hydrothermal fluids are strongly alkaline, a second liquid-liquid extraction was 203 performed to extract any organic acid anions that may have been present following acidification 204 to pH < 3 by addition of concentrated HCl. The $pH_{25^{\circ}C}$ values of the Rainbow and Lucky Strike 205 fluids were well below the pK_a of alkanoic acids, so these samples were not acidified.

206 The DCM extracts were concentrated by evaporation under a gentle stream of N₂ at room temperature. When the extract had been concentrated to ~ 2 ml, it was capped and transferred to 207 208 a freezer at -20°C, which in most cases resulted in exsolution of residual water to form a ring of 209 ice within the sample vial at the top of the solvent. When this occurred, the solvent phase was 210 transferred to a new vial, concentrated further, and then stored at -20° C until analysis. During 211 concentration of these and all other extracts, care was taken to prevent the solvent from 212 evaporating to dryness to limit the loss of semi-volatile organic compounds. Prior to analysis, 213 the concentrated extracts were treated with Cu powder to remove native sulfur. The Cu for this 214 treatment was activated by soaking in concentrated HCl and then rinsed with organic solvents.

215 After removal of the hydrothermal fluid, the interior of the IGT sampler was rinsed with 216 successive aliquots of methanol (MeOH) and DCM to collect organic compounds that may have 217 precipitated on the walls of the sampler. Both rinses were combined into a single glass vial. In 218 practice, a small amount of residual hydrothermal fluid remained in the sampler that was 219 collected along with the solvents. The combined solvents were evaporated to approximately half 220 their original volume with a stream of N₂, at which point the extract usually separated into an 221 H₂O/MeOH-rich phase underlain by a DCM-rich phase. Additional DCM and acidified Fisher 222 H₂O were added to the vial to enhance transfer of organic compounds to the DCM phase. To 223 reduce the possibility of contaminating the samples with organic compounds, the H₂O used for 224 this step was extracted with DCM before adding it to the solvent extract. After shaking, the 225 DCM-rich layer was pipetted into a separate vial and processed using the same methods 226 described for the fluid extract.

227 To extract organic compounds precipitated on the walls of the Ti tubing of the POP gun, 228 remaining fluid was removed by suction with a syringe and the tubing was then rinsed with ~5 229 ml aliquots of MeOH and DCM in succession. The solvents were pushed through the tubing 230 using a solvent-cleaned glass syringe with Teflon plunger. The MeOH used for this procedure 231 was mildly acidified with HCl to ensure extraction of carboxylic acids and other weakly acidic 232 compounds present as anions. In some cases, this had the unforeseen consequence of converting 233 organic acids to methyl esters. While most of the MeOH extracts from the POP gun samples 234 were found to contain small amounts of organic compounds, very few of the DCM extracts 235 contained detectable organic compounds, indicating that most soluble organic compounds were 236 removed from the tubing by the MeOH rinse.

237 Because the MeOH extracts from the POP gun inevitably contained a small amount of 238 hydrothermal fluid that remained in the tubing, the extracts were partially concentrated under a 239 stream of N₂, and DCM added. The combined MeOH-DCM extract was then placed in a -20 °C 240 freezer to induce phase separation, and the DCM-rich (lower) phase was removed to a clean vial 241 for further processing (note that although DCM was added to these samples, they are still 242 referred to as "MeOH extracts" to reflect the solvent used in the initial extraction step). These 243 extracts as well as the original DCM rinses from the POP gun samples were further concentrated 244 using the same methods described for the IGT fluid samples.

To process the SPE cartridges, hydrothermal fluids remaining in the cartridges was first removed by suction using a large, gas-tight syringe. A glass/Teflon syringe was then used to pass ~10 ml aliquots of MeOH and DCM through the cartridges in succession. To ensure extraction of organic acid anions, the solvents for this step were mildly acidified with HCl. The MeOH and DCM rinses were collected in separate pre-cleaned glass vials and processed separately using the methods described above for the IGT and POP gun extracts.

251 In addition to the fluid samples, organic compounds were extracted from a portion of carbonate chimney from Lost City. The extracted chimney sample was a large fragment (36 g, 252 253 wet weight) of a small, active flange collected from the Beehive/EXOMAR-12 site adjacent to 254 where the vent fluid samples were collected. Because the flange crumbled into pieces during 255 collection by the ROV, it is not possible to precisely determine what part of the structure the 256 fragment represents, but it was mostly composed of material from the flange interior. The 257 chimney fragment was crushed in a slurry of 80% DCM/20% MeOH, transferred to a pre-258 cleaned glass bottle, and sonicated twice for 10 minutes with vigorous shaking of the sample in 259 between. Following removal of the extract to a separate bottle, the chimney was extracted two 260 additional times using 80% DCM/20% MeOH followed by 100% DCM. When all of the 261 extracts were combined, the solvents separated into two phases, presumably because of the 262 presence of seawater or vent fluid in the chimney sample. At this stage, the solvent was acidified 263 with HCl to ensure that alkanoic acids would enter the organic solvent, and the lower, DCM-rich 264 phase was then removed to a separate bottle. In order to remove inorganic salts, the solvent extract was rinsed twice with Fisher DI water that had been pre-extracted with DCM to remove 265 organic compounds. The resulting solvent extract was concentrated under a stream of N₂ and 266 prepared for analysis using procedures similar to those used for the fluid samples. 267

268 **2.3. Analytical methods**

269The solvent extracts were analyzed onshore using gas chromatography-mass spectroscopy270(GC-MS) on an Agilent 6890 interfaced to a 5973 mass selective detector. Routine analyses271were performed with an Alltech AT-5ms column (60 m length, 320 µm internal diameter, 0.25272µm film thickness) using a temperature program of 50 °C initial temperature for 5 min, 10

²⁷³ °C/min to 325 °C, and hold at this temperature for 30 min. For chimney samples, organic acids ²⁷⁴ in an aliquot of the solvent extract were converted to fatty acid methyl esters (FAME) prior to ²⁷⁵ analysis by reaction with BF_3 in methanol using kits from Supelco. Compounds were routinely ²⁷⁶ identified by comparison of mass fragment spectra with those in the NIST02 reference library. ²⁷⁷ In a few cases where the spectral identifications were ambiguous (e.g., phenanthrenes, ²⁷⁸ hexadecanol), retention times and mass spectra were confirmed by comparison with analyses of ²⁷⁹ known compounds.

280 During analysis of extracts by GC-MS, a number of compounds were observed at short 281 retention times (<24 min) that could be attributed to trace contaminants from the solvents which 282 had been concentrated in the extracts during sample preparation (note that these relatively 283 volatile compounds were lost in some samples that were inadvertently evaporated to dryness). 284 Prominent among these were undecane and several methyldecane isomers that was found to be a 285 trace contaminant in the MeOH solvent, and several small methylcycloalkanes present in the 286 DCM. Owing to interference from these compounds, low molecular weight organic compounds 287 that may have been present in the hydrothermal fluid samples with short retention times could 288 not be confidently identified during this study. In addition, the sorbents used in the SPE 289 cartridges were found to release relatively large amounts of a spectrum of organic compounds 290 during sample processing, most of which did not match any compounds in the reference library 291 and could not be identified. These contaminants are discussed more fully below, and example 292 mass fragmentograms for the unknown compounds are provided as Supplemental Figure S1.

293 Compound specific carbon isotope analyses were performed on select samples with an 294 Agilent 6890 coupled to a Delta plus XL mass spectrometer via a combustion interface operated 295 at 1080 °C using a column and temperature program similar to the GC-MS analyses. Instrument 296 error was estimated using an external standard containing 15 *n*-alkanes with known δ^{13} C values 297 and was measured to be 0.2‰ during the period when these samples were analyzed. All extracts 298 were derivatized using BSTFA +1% TMCS and pyridine and heated at 70 °C for one hour. 299 Results were corrected for the addition of methyl carbon by derivatizing phthalic acid with a 300 known isotopic value using the same BSTFA and subtracting the contribution of methyl carbons 301 by mass balance.

302 Quantification of fatty acids in selected samples was performed by conversion of the acids to 303 methyl esters and comparison of peak areas for reconstructed ion chromatograms with reference 304 standards for octanoic and dodecanoic acids. Analysis of hydrothermal fluid samples for the presence of dissolved free amino acids was performed by gas chromatography with flame ionization 305 306 detection following extraction and derivatization of the amino acids with EZ:faast kits from 307 Phenomenex Inc. (Torrance, CA), using the column and analytical parameters specified by the 308 manufacturer. Although an effort was made to detect highly polar compounds by direct injection of 309 small aliquots (1-2 µl) of hydrothermal fluid into the GC-MS using an AquaWax column (Alltech), 310 no additional organic compounds were identified by this method.

311 **2.4.** Assessment of potential sources of background contaminants

In an effort to evaluate potential exogenous sources of organic compounds that might contaminate the hydrothermal fluid samples, samples of bottom seawater were extracted and analyzed using the same procedures utilized for the hydrothermal samples (Fig. 4). A sample of bottom seawater was obtained with the IGT sampler in the vicinity of the TAG field (J2-364-IGT6), but no organic compounds were found in either the DCM extract of the fluid or in the solvent rinse of the sampler. Similarly, no organic compounds were observed in a DCM extract of the Fisher H_2O used in sample processing. Samples of bottom seawater were obtained with 319 the POP gun in the vicinity of the Lucky Strike and Lost City vent fields (J2-356-POP1, J2-362-320 POP1). The MeOH extracts of both seawater samples exhibited a large, unresolved complex 321 mixture (UCM) of organic compounds that were dominated by aliphatic components (Fig. 4a; 322 see Supplemental Fig. S2 for an example mass fragmentogram of the UCM). The UCM 323 displayed three distinct peaks centered at retention times of ~26, ~29.5 and ~36 min. Analyses 324 of the oils used for pressure compensation and hydraulic parts on the Jason submersible are also 325 characterized by broad UCM humps (Figs. 4b and 4c). However, the UCM humps for the oils 326 show very different distribution patterns than those observed for the seawater samples, indicating 327 that the UCM observed in the seawater samples included little or no contribution from the 328 submersible. No organic compounds were found in the DCM extracts of the seawater POP gun 329 samples.

330 Solvent extracts of the SPE cartridges that were attached to the POP gun during sampling of 331 bottom seawater [(J2-356-SPE(HLB), J2-362-SPE(C8)] contained numerous organic 332 compounds, most of which did not match any compounds in the spectral library used in the GC-333 MS analysis (Figs. 4d and 4e). Similar suites of compounds were observed in all other samples 334 processed through the SPE cartridges, including the hydrothermal fluids and 1 L samples of 335 Fisher H₂O pumped through the cartridges and processed using the same procedures. 336 Consequently, the compounds seen in the extracts of the seawater samples appear to be derived 337 predominantly from the HLB and C8 SPE phases rather than the fluid samples, and likely 338 represent fragments of the polymers used to coat the extraction beads. In the case of the C8 339 phase, the suite of compounds seen in the extracts for all samples included small amounts of C_{16} 340 and C_{18} alkenes and alkenones, which also appeared to be derived from the SPE beads. While 341 trace amounts of organic compounds derived from seawater may be present in the SPE samples, 342 they could not be distinguished from the pervasive background of compounds from the SPE 343 phases.

344 3. RESULTS

A summary of the organic compounds that were positively identified in the hydrothermal vent samples is provided in Table 2, with more detailed discussion in the following sections. Note that some compounds in this table were identified using methods that were not applied at all sites, so results for different sites may not be directly comparable. For brevity, only the most salient results of the GC-MS analyses are shown in the figures displayed here; figures showing additional analyses of samples are provided as Supplemental Materials.

351 The concentration of Mg in fluid samples collected with the IGT samplers are listed in Table 352 1. Magnesium concentrations of fluids sampled from deep-sea hydrothermal vents are routinely 353 used to estimate the amount of seawater entrained into the fluid during sampling, based on the 354 presumption that the vent fluids have Mg concentration approaching zero as a result of fluid-rock 355 interactions in the subsurface while seawater has a concentration of ~52.2 mmol/kg (German and 356 Seyfried, 2014). However, analysis of IGT samples obtained for this study indicates that 357 endmember vent fluids at the Rainbow field may have Mg concentrations of 1.5-2 mmol/kg (see Seyfried et al., 2011). Only two of the samples collected with the IGT samplers had elevated Mg 358 359 contents (US4, Kremlin) indicating a substantial seawater component. The remaining samples 360 have Mg concentrations < 8 mmol/kg and most have concentrations of 2.6 mmol/kg or less, 361 indicating nearly pure hydrothermal fluids were obtained. Also listed in Table 1 for reference 362 are measured pH values (at 25 °C) and endmember concentrations of dissolved H₂, CO₂, and

363 CH₄ in the fluids determined from separate IGT samples obtained at the same chimney orifices 364 as the organic samples (Reeves et al., 2014).

365 **3.1. Lost City**

366 Lost City was the only vent field where significant amounts of extractable organic compounds were consistently observed in the hydrothermal fluid samples. The initial liquid-367 liquid extracts of hydrothermal fluids from the Beehive site at Lost City collected with the IGT 368 369 samplers yielded no detectable organic compounds (Fig. 5a), but several compounds were 370 present in the sample from the top of Poseidon, including trithiolane, hexadecanol, and squalene 371 (Fig. 5e). Furthermore, hexadecanol, squalene, and cholesterol were observed in the solvent rinse 372 of the IGT sampler used to collect fluids at the top of Poseidon (Fig. 5f), indicating that a 373 fraction of these compounds may have precipitated onto the walls of the sampler. No organic 374 compounds were observed in solvent rinses of the IGT samplers for the Beehive site.

375 Following acidification and re-extraction of the Beehive and Poseidon fluids, a suite of three 376 alkanoic acids was observed that included *n*-octanoic, *n*-decanoic and *n*-dodecanoic acids (C_8 , C₁₀, and C₁₂, respectively) (Fig. 5b & 5f). These compounds were particularly abundant in the 377 samples from the Beehive site (Fig. 5b), with relatively low levels observed in the sample from 378 379 the top of Poseidon (Fig. 5f). Both fluid samples from Beehive yielded similar alkanoic acid 380 concentrations that decreased with increasing carbon number ($C_8 > C_{10} > C_{12}$) from values of 0.72 to 0.23 µmol/L (Table 3). In both samples, acid concentrations decreased with increasing 381 382 carbon number ($C_8 > C_{10} > C_{12}$). Since it is not clear that the extraction methods completely 383 removed the acids from these fluid samples, the reported concentrations should probably be 384 regarded as minimum values. The alkanoic acids are characterized by relatively heavy carbon isotope compositions, with values between -2.8‰ and -7.6‰ (Table 3). The extract of the 385 386 acidified fluid for the Poseidon sample also included trithiolane, hexadecanol, and squalene as 387 well as trace amounts of cholesterol, indicating that the initial liquid-liquid extraction prior to 388 acidification did not completely remove these compounds (Fig. 5f).

389 Conspicuously absent from the Lost City IGT sample extracts were alkanoic acids with fewer than 8 or greater than 12 carbon atoms (e.g., C_6 , C_{14}). To evaluate whether the absence of these 390 391 compounds may have been an artifact of the analytical protocols, an artificial seawater solution 392 with pH adjusted to 9.5 containing ~200 μ g/L each of C₆, C₁₂, and C₁₄ alkanoic acids was 393 prepared and processed with the same methods used for the natural samples. Quantitative analyses of the resulting extracts showed that recovery of the C14 acid was essentially identical to 394 395 the C₁₂ acid, implying that the absence of the C₁₄ acid in the Lost City extracts accurately reflects 396 its abundance in these samples. On the other hand, recovery of the C₆ acid was only about 40% 397 relative to the C_{12} acid, perhaps because extraction of the C_6 acid from the water phase by the 398 solvent was less efficient owing to its relatively higher polarity or because this compound was 399 partially volatilized during sample concentration. Nevertheless, even partial recovery of the C₆ 400 acid suggests that it would have been detected in the samples if it had been present at levels 401 comparable to the other acids. Accordingly, it appears that levels of both the C₆ and C₁₄ acids in the Lost City hydrothermal fluids were significantly lower than those of the C_8-C_{12} acids. 402

403 The two POP gun samples at Lost City (J2-361-POP1, J2-361-POP2) were both obtained at 404 the Beehive site, and the MeOH extracts were found to contain C_{10} and C_{12} alkanoic acids 405 similar to those observed in the IGT extracts from this location (Fig. 5c). However, the relative 406 proportions of these compounds were reversed from the IGT samples, with the C_{10} acid present 407 in much lower amounts than C_{12} . The C_8 alkanoic acid was not observed in either MeOH extract 408 of the POP gun samples. No organic compounds were detected in DCM extracts of the POP409 guns.

410 Despite the presence of contaminants from the extraction phase, organic compounds that 411 could be confidently attributed to the hydrothermal fluid were identified in the Beehive sample 412 that was pumped through the HLB SPE phase [J2-361-SPE(HLB)] and extracted with MeOH 413 (Fig. 5d). This sample contained large amounts of the same suite of C_8 , C_{10} , and C_{12} *n*-alkanoic 414 acids that were seen in the IGT samples, and these compounds were present at levels far above 415 those of any background contaminants. Estimates of the concentrations of the acids in the 416 hydrothermal fluid were calculated from the measured abundance of the compounds in the 417 extract and the amount of fluid pumped through the SPE cartridge (~1.2 L). The resulting concentrations are similar to those measured for the Beehive IGT samples, and show the same 418 419 decrease in concentration with increasing carbon number (Table 3). No organic compounds 420 were identified in the DCM extract of the HLB phase other than those that could be attributed to 421 background sources. Fluids from the Beehive site were also pumped through a cartridge 422 containing the C8 SPE microbeads [J2-361-SPE(C8)]. However, no identifiable compounds that 423 could be attributed to the hydrothermal fluid were found in solvent extracts of this phase (see 424 Supplemental Fig. S3).

425 To aid in evaluation of potential sources for alkanoic acids observed in the Lost City vent 426 fluids, organic compounds were extracted and analyzed from a piece of carbonate chimney 427 collected from the Beehive site where the IGT and POP gun samples were obtained. As shown 428 in Figure 6a, analysis of the total extract by GC-MS was dominated by a UCM, although several 429 individual hydrocarbons could be identified including *n*-hexadecane and squalene. Also 430 prominent in the total extract are a number of wax esters (Fig. 6a). These compounds apparently 431 formed through condensation of alkanoic acids with alkanols and alkenols in the extracts induced 432 by acidification of the solvents.

433 In order to determine the distribution and diversity of alkanoic acids in the chimney, the acids 434 in an aliquot of the total extract were converted to fatty acid methyl esters (FAME) (Figs. 6b and 435 6c). The FAME analysis revealed an assortment of alkanoic and alkenoic acids ranging in 436 carbon number from C₈ to C₂₈, with a strong predominance for compounds with an even number 437 of carbon atoms. The C₁₄, C₁₆, and C₁₈ compounds are substantially more abundant than 438 compounds with other numbers of carbon atoms, and occur in approximately equal proportions. 439 There is a sharp drop in abundance of the alkanoic acids with fewer than 14 carbon atoms. 440 While it is possible that this drop off could be partially explained by loss of the relatively volatile 441 shorter-chain acids during concentration of the sample, the same procedures used to minimize 442 loss of semi-volatile compounds in the hydrothermal fluid extracts were employed in sample 443 preparation of the chimney extract, suggesting that the relative amounts of these compounds are likely to reflect their in situ abundance. The C₁₄-C₁₈ carboxylic acids in the chimney extract had 444 445 carbon isotopic compositions between -13.5% and -20%, substantially lighter than the acids in 446 the fluid (see Supplemental Table S1 for a complete listing of isotopic analyses for the chimney 447 extract).

448 **3.2. Lucky Strike and Rainbow**

449 Most solvent extracts of the fluids collected at Lucky Strike and Rainbow using the IGT 450 samplers were found to contain either no detectable organic compounds or trace amounts of 451 compounds that appeared to be mostly derived from sources other than the hydrothermal fluid 452 (Fig. 7). Several of the IGT extracts contained small peaks whose mass spectra did not provide a

453 close match to any compounds in the spectral library, and were inferred to be contaminants, 454 possibly fragments of plastic polymers (see Supplemental Fig. S1 for an example). The source 455 of these compounds could not be determined. Several of the fluid extracts were also found to 456 contain small amounts of Fluorolube lubricant, but these compounds could be readily identified 457 as contaminants by their mass fragmentation patterns during GC-MS analysis. An example is 458 provided in Figure 7b, which shows the extract for a fluid sample from the Lucky Strike 459 hydrothermal field (J2-359-IGT6). The only identifiable compound in this sample that could not 460 be attributed to Fluorolube or other contaminants was a small peak for dibenzothiophene. 461 Dibenzothiophene was present in one other sample, also from Lucky Strike (J2-358-IGT6), but 462 was not found in any other samples analyzed for this study. Analyses of the solvent rinses of the 463 IGT samplers following removal of the fluid did not reveal the presence of any compounds that 464 could be confidently attributed to a hydrothermal fluid source. Most of these rinse samples, 465 however, did include small amounts of phthalates from an unidentified source as well as traces of 466 Fluorolube.

467 In contrast to the scarcity of organic compounds in the IGT samples, organic compounds 468 were observed in several of the POP gun samples from sites at Rainbow and Lucky Strike. 469 Extracts from all sites at Rainbow sampled with the POP gun contained small amounts of PAH 470 that included phenanthrene, methylphenanthrenes, and C₂-phenanthrenes (e.g., Figs. 8a-c). In 471 addition, samples from the Guillaume vent contained hexadecanol and a C₁₈ alkenoic acid (Fig. 472 8c). Because the solvents used in the extracts were mildly acidified, the organic acids in this and 473 other samples were converted to methyl esters during sample processing. While it is possible 474 that these compounds represent membrane components of microorganisms entrained into the 475 POP gun during sampling, essentially identical results were obtained from both POP gun 476 samples at this site, indicating it is more likely that the compounds were indigenous to the 477 hydrothermal fluid.

478 Extracts of the POP gun samples from the high temperature Stylo 2 and Padraig sites at 479 Rainbow also contained a small, narrow UCM hump centered at a retention time of ~36 min 480 (Fig. 8b). A UCM hump was also evident in both samples from the lower temperature (191 °C) 481 Ecurie site at Rainbow, but in this case the UCM was much broader and extended from 26 to 38 482 min retention time (Fig. 8a). In all cases, the UCM was dominated by aliphatic components as 483 indicated by a predominance of mass fragments with mass/charge ratios (m/z) of 55, 57, 69, 71, 484 83, and 85 (see Supplemental Fig. S2). The single POP gun sample obtained at Lucky Strike did 485 not contain a prominent UCM, but did include several even-carbon-numbered alkenoic and 486 alkanoic acids ranging in carbon number from C_{14} to C_{22} , with C_{16} and C_{18} alkenoic acids 487 particularly abundant (Fig. 8d). Since these compounds are predominant components of 488 bacterial membranes, it appears likely that in this case the alkenoic and alkanoic acids must have 489 been extracted from biomass entrained into the POP gun during sampling rather than being 490 dissolved in the hydrothermal fluid.

491 Among the DCM extracts of the POP gun samples, organic compounds were observed in 492 only some of the Rainbow samples and were not found in any samples from Lucky Strike. 493 Extracts of samples from the high temperature Guillaume and Padraig sites at Rainbow all 494 contained phenanthrene and methylfluorenes present in trace amounts that were just barely 495 detectable (Fig. 9a). However, these compounds were not observed at the other high temperature 496 site sampled, Stylo 1. One of the DCM extracts for the POP gun samples from the lower 497 temperature Ecurie site at Rainbow contained a homologous series of *n*-alkanes ranging in 498 carbon number from 23 to 33 (Fig. 9b). The alkanes exhibited a bell-shaped distribution in

499 abundance centered around $\sim C_{28}$, with no apparent even- or odd-carbon-number preference. 500 These hydrocarbons were not observed in any other POP gun extract.

None of the MeOH extracts of the HLB cartridges from the Rainbow and Lucky Strike systems contained organic compounds that could be differentiated from the background of compounds derived from the SPE sorbent (see Supplemental Fig. S4). The same suite of background peaks observed in the MeOH extracts were also evident in the DCM extracts of the HLB phase (Supplemental Fig. S4). In addition, the DCM extracts also included C_{16} and C_{18} alkenes that were not present in the MeOH extracts, but which also appeared to be derived from the extraction phase since they were present in all samples processed through the SPE cartridges.

508 Despite the ubiquitous presence of background contaminants, organic compounds that could 509 be attributed to the hydrothermal fluids were identified in some DCM extracts of the HLB phase. 510 The DCM extract of sample J2-359-SPE(HLB) from the Medea site at Lucky Strike (Fig. 10b) 511 was found to contain methyl esters of C₁₆ and C₁₈ alkanoic and alkenoic acids similar to those 512 observed in the MeOH extract of the POP gun at this same sample location (Fig. 8d). The DCM 513 extract of the HLB phase obtained from the Ecurie site at Rainbow (Fig. 10a) contained a 514 homologous series of C23-C33 n-alkanes with a distribution identical to that observed in the DCM extract of the POP gun sample for this site (Fig. 9b). In both cases, the compounds identified 515 516 were not observed in DCM extracts of the SPE cartridges from other sites.

The only high-temperature sample processed through a SPE cartridge containing the C8 sorbent was from the Stylo 2 site at Rainbow [J2-355-SPE(C8)]. Extracts of this sample contained the same suite of compounds observed in all other C8 extracts that appeared to be attributable to background from the sorbent phase (Supplemental Fig. S3). Although there is a broad "window" during the GC-MS analysis between 16 and 30 min retention time where there is little or no background, no organic compounds indigenous to the sample were observed in this interval.

524 **3.3. Amino acid analyses**

525 Preserved fluid samples from the IGT samples from all sites were analyzed for the presence 526 of free amino acids by GC-MS following extraction from the fluid and derivatization. However, no free amino acids were detected in any of the hydrothermal fluids or in bottom seawater 527 samples at a detection limit of 5 nmol kg⁻¹. This result is consistent with the absence of 528 detectable free amino acids at similar levels in 319 °C hydrothermal vent fluids from the 529 530 Guaymas Basin (Haberstroh and Karl, 1989) and >200 °C hydrothermal vent fluids from the Marianas Trough (Fuchida et al., 2014). Conversely, Klevenz et al. (2010) reported several 531 dissolved amino acids present above the 5 nmol kg⁻¹ threshold from high-temperature 532 hydrothermal sites further south on the Mid-Atlantic Ridge, with total free amino acid 533 concentrations up to 377 nmol kg⁻¹. Lang et al. (2013) reported total hydrolysable amino acids 534 (THAA) for hydrothermal fluids from the Lost City system ranging from 736 to 2300 nmol kg⁻¹, 535 536 but did not report free amino acid concentrations. The apparent absence of detectable free amino 537 acids in the Lost City fluids suggests that the amino acids reported by Lang et al. (2013) are 538 predominantly present as peptides or proteins.

539 4. DISCUSSION

540 **4.1. Organic compounds in Lost City hydrothermal fluids**

541 Lost City was the only sampling location in this study where significant amounts of dissolved 542 organic compounds were consistently observed in the hydrothermal fluids. At that site,

543 substantial concentrations of C8-C12 alkanoic acids as well as several other organic compounds (squalene, hexadecanol, trithiolane, and cholesterol) were observed in multiple samples and, in 544 545 the case of the alkanoic acids, were identified in all samples regardless of the collection and 546 processing methods used (Fig. 5). Although several of the compounds found in the Lost City 547 fluids were also present at low abundance in extracts of the carbonate chimney recovered from 548 that site, the fluid extracts lacked numerous other compounds that were abundantly present in the 549 chimney extract. We therefore conclude that the organic compounds observed in the Lost City 550 samples were indigenous to the hydrothermal fluid and not derived from chimney particles 551 entrained during sampling.

552 Measured concentrations of the C8-C12 acids for the two IGT samples and the SPE sample at 553 the Beehive are in good agreement, and indicate concentrations for individual acids in the 0.23 to 554 0.77 µmol/L range, which is equivalent to 2.7-6.2 µmol C/L (Table 3). For comparison, 555 concentrations of dissolved formate and acetate for fluids venting at the Beehive site in 2005 were found to be ~144 μ mol/L and 8.5 μ mol/L, respectively (Lang et al., 2010). The measured 556 557 concentration of dissolved organic carbon (DOC) in the fluids sampled in 2005 was $\sim 102 \mu$ mol 558 C/L, of which \sim 50 µmol C/L can be accounted for as formate and acetate (note that some of the 559 formate and acetate are lost during sample preparation for DOC analysis) (Lang et al., 2010). 560 Collectively, the C₈-C₁₂ acids can account for ~11-15 µmol C/L, or 11-15% of the DOC (Table 561 3). Taken together, it appears that the carboxylic acids that have been measured so far (including 562 formate) can account for >60% of the DOC in the hydrothermal fluids at the Beehive site. Total hydrolysable amino acids can account for an additional 3-5% of DOC (Lang et al., 2013), 563 564 leaving about one third of the total DOC still unaccounted for at this point. A number of 565 alkylated monocyclic compounds have been identified in the Lost City fluids using other 566 methods (Konn et al., 2009) that might contribute to the remaining third, or it could be composed 567 of highly polar molecules (e.g., pyruvate, alkanols) not detectable by the methods employed to 568 date. Alternatively, the sampling methods used may have underestimated the amounts of 569 carboxylic acids present.

570 The detection of only even-carbon-numbered alkanoic acids in the Lost City samples 571 indicates that the acids have a biological source. Although abiotic organic synthesis reactions 572 such as the Fischer-Tropsch process are known to produce long-chain carboxylic acids (e.g., 573 McCollom et al., 1999, 2010), these reactions typically produce an mixture of even- and odd-574 carbon-numbered compounds in roughly equal amounts as the result of sequential addition of 575 single carbon units (McCollom and Seewald, 2007). There is currently no known mechanism for 576 abiotic synthesis of organic compounds in two-carbon units that might produce exclusively even-577 carbon-numbered compounds, nor is there any evidence from natural organic products thought to 578 have an abiotic abiotic origin for an even-number predominance. For instance, abiotic organic 579 compounds in meteorites have approximately equal proportions of even- and odd-carbon-580 numbers (Yuen and Kvenvolden, 1973; Huang et al., 2005). Furthermore, the absence of other 581 types of even-carbon-numbered compounds (e.g., hydrocarbons) makes it unlikely that the acids 582 come from an abiotic source.

Since the C_8 - C_{12} alkanoic acids are too small to be components of lipid membranes and are not known as common metabolic products of microbial metabolism, the biological source for the compounds is unclear. However, these compounds are known to be inhibitors of some metabolic pathways such as fermentation (e.g. Legras et al., 2010), and perhaps they play a similar role in the Lost City microbial community. These compounds are probably produced by bacteria, since archaea do not appear to synthesize linear alkanoic acids. Studies of the microbial communities

at the hottest Lost City vents (70-91 °C) have found that the interiors of the carbonate chimneys 589 590 are dominated by a single species of archaea related to the Methanosarcinales, but the chimney 591 exteriors and hydrothermal fluids contain bacteria that could be the source of the alkanoic acids 592 (Schrenk et al., 2004; Brazelton et al., 2006). Alternatively, the compounds might be produced 593 in the subsurface. In either case, the presence of C_8 - C_{12} alkanoic acids at relatively high 594 abundance may reflect novel metabolic pathways taking place in this high pH ecosystem. 595 Another possibility is that the compounds are formed through thermal decomposition of 596 bioorganic matter in the subsurface. In this respect, it may be notable that recent studies of 597 insoluble carbonaceous matter from subseafloor serpentinites have found that they contain a 598 large component of aliphatic compounds up to C_{12} associated with carboxylate functional groups 599 (Pasini et al., 2013).

600 The alkanoic acids in the fluid and chimney samples from Lost City form complementary 601 sets, with the fluid containing C_8 , C_{10} , and C_{12} acids but none of the longer chain acids, while the chimney contained substantial amounts of C_{14} , C_{16} , and C_{18} acids with only small amounts of the 602 603 shorter chain acids (Figs. 5 and 6). This result suggests two possible origins for the acids 604 dissolved in the Lost City hydrothermal fluids. The first possibility is that a complete suite of C_8 to C_{18} acids is produced by biological processes occurring within the chimney walls, but the 605 shorter acids are preferentially dissolved by the circulating hydrothermal fluids owing to their 606 607 relatively higher aqueous solubility. Alternatively, all of the acids may be produced by biological activity deeper in the system that are mobilized by the hydrothermal fluids, but only 608 609 the longer-chain acids are subsequently removed by precipitation as the fluids cool during 610 mixing with seawater within the walls of the chimney (a mechanism that would be directly 611 comparable to that used as the basis for the design of our POP gun sampling device). In the 612 former case, mobilization of the shorter acids would represent a significant loss of fixed organic 613 carbon from the microbial communities living in the chimney walls, while the latter case would 614 represent a significant net transfer of organic carbon from deep within the system to the seafloor 615 chimney environment.

616 The carbon isotopic compositions of the C_8 - C_{12} alkanoic acids do not allow for clear 617 discrimination between these two possibilities. The C_8 - C_{12} acids in the fluid are relatively heavy compared to the C_{14} - C_{18} acids in the adjacent chimney (-2.8% to -7.6% vs. -13.5% to -20%) 618 (Fig. 11). However, the C_{16} and C_{18} acids in other chimney samples from Lost City analyzed by 619 620 Bradley et al. (2009) overlap with the heavier values observed in the fluids, with some values as low as -1.1% (Fig. 11). Thus, it appears possible, and maybe even likely, that the C₁₄-C₁₈ acids 621 in the Lost City chimneys are composed of a mixture of heavier compounds from the same 622 623 source as the acids observed in the hydrothermal fluid and lighter compounds derived from in 624 situ microbial activity, with the relative contributions from these sources varying locally. Within 625 this scenario, however, the ultimate source of the heavier acids could either be local to the 626 chimney deposits or somewhere deeper within the system.

627 The IGT fluid sample collected at the top of the Poseidon structure contained substantially 628 lower amounts of the C₈-C₁₂ alkanoic acids than the samples collected from the Beehive site on 629 the side of Poseidon (Fig. 5), and also contained a number of compounds not observed in the 630 other samples including squalene, trithiolane, hexadecanol, and cholesterol. The fluid from the 631 Poseidon summit also had a substantially lower temperature (51 °C vs. 91 °C), indicating the 632 fluid had conductively cooled or mixed with cold seawater during passage through the length of 633 the structure. Squalene is a common product of the microbial community living within Lost City 634 chimney structures (Méhay et al., 2013), and microbial activity within the chimney structure is

635 presumably the source of this compound as well as the hexadecanol and cholesterol. Although 636 O₂ is required for biosynthesis of cholesterol, influx of seawater into chimney mixing zones could supply O_2 for aerobic organisms, and the cholesterol was likely entrained into circulating 637 638 fluids from these zones. Microbial or other abiotic processes occurring within the chimney may 639 have also led to partial loss of alkanoic acids in the fluid. To our knowledge, trithiolane has not 640 previously been reported to occur in deep-sea hydrothermal systems. However, we have 641 observed trithiolane and several other cyclic carbon-sulfur compounds in strongly alkaline (pH >642 11) fluids discharged from serpentinite-hosted springs at Aqua de Ney in northern California 643 (Feth et al., 1961; Barnes et al., 1972), suggesting these compounds may be common in fluids discharged from serpentinites (Fig. 5g). Whether the trithiolane observed in the Lost City fluid 644 645 is a biological or abiotic product is presently unclear.

646 **4.2. Organic compounds in high-temperature hydrothermal fluids**

647 The high-temperature (\geq 190 °C) hydrothermal fluids examined for this study were 648 characterized by an overall scarcity of extractable dissolved organic compounds. Most samples 649 were found to contain either no detectable organic compounds or only trace amounts of organic 650 compounds composed primarily of polycyclic aromatic hydrocarbons. More substantial amounts 651 of organic compounds were observed in a few of the POP gun and SPE samples but, as discussed 652 further below, these compounds appeared to represent particulate matter or biomass entrained 653 during sample collection rather than dissolved components present in the hydrothermal fluids.

654 Several of the MeOH extracts from the POP gun samples at Rainbow were found to contain unresolved complex mixtures of organic compounds (e.g., Figs. 8a and 8b). The UCMs 655 observed for the hydrothermal samples have some similarities to the UCM found in the POP gun 656 657 samples of deep seawater, but there are also some substantial differences. The UCM observed 658 for bottom seawater displayed a broad hump extending from ~ 25 to ~ 37 min retention time, with 659 three distinct peaks centered at ~26, ~29.5 and ~36 min (Fig. 4a). The UCM for the moderately 660 hot fluids (191 °C) from the Ecurie site also displayed a broad hump extending from ~25 to ~37 661 min (Fig. 8a). However, while the Ecurie samples display a distinct peak at 36 min similar to that for seawater, there is only a slight indication of a peak at 29.5 min and none at 26 min. 662 Conversely, the UCM humps evident in the samples from the \geq 350 °C Stylo 2 and Padraig sites 663 are much narrower, with a single peak at 36 min that coincides closely with the final peak in the 664 665 seawater UCM (Fig. 8b). For both seawater and hydrothermal fluid samples, mass fragmentation patterns from GC-MS analyses show that the UCMs are dominated by saturated alkyl (m/z = 57, 666 667 71, 85, etc.) and monounsaturated alkenyl (m/z = 55, 69, 83, etc.) fragments, indicating the compounds that comprise the UCMs have a large aliphatic component (see Supplemental Fig. 668 669 S2). This observation appears to be consistent with studies indicating that dissolved organic 670 matter (DOM) in deep Atlantic Ocean water contains a large aliphatic component (Hertkorn et 671 al., 2013), although the actual molecular structures of the compounds making up the UCM 672 remain to be determined.

673 The similarities between the hydrothermal and seawater UCMs suggest that the mixture of compounds that comprise the hydrothermal UCMs may represent seawater organic matter that 674 675 has been modified by high-temperature processes or biological activity during circulation though 676 the hydrothermal system. If this is the case, the hydrothermal samples appear to have 677 preferentially lost the lower molecular weight fraction of the UCM, with more extensive loss occurring at the higher temperature Stylo 2 and Padraig sites than at the lower temperature 678 679 Ecurie site. Regardless, differences between the seawater and hydrothermal UCMs indicate that 680 the organic compounds found in the hydrothermal samples do not simply represent seawater

organic matter entrained during sample processing or shallow mixing of seawater into the fluids.
 In contrast, the fluids from Lucky Strike and Lost City contained no apparent UCM, suggesting
 that the compounds that constitute seawater UCM are quantitatively removed during circulation
 through these systems.

685 Although the UCMs observed in the POP gun samples from Ecurie represent the largest 686 amount of organic matter observed in any of the high-temperature fluid samples in this study, 687 these compounds appear to be derived from particulate materials entrained during sampling 688 rather than compounds dissolved in the hydrothermal fluid. Extracts of hydrothermal chimneys 689 from the Rainbow field also contain substantial UCM humps (Simoneit et al., 2004). While 690 direct comparisons are complicated by differences in analytical protocols, the UCM humps for 691 many chimney samples reported by Simoneit et al. (2004) are broad with a peak towards the 692 higher molecular weight fraction, and appear to be very similar to the UCMs observed for the 693 Ecurie samples. In addition, many of the chimney samples display a suite of linear alkanes ranging in carbon number from $\sim C_{22}$ to $\sim C_{33}$. This suite of alkanes displays a bell-shaped 694 distribution in abundance centered at about C₂₈, with no even-odd carbon number preference. 695 696 The carbon number range and abundance distribution of these compounds is nearly identical to 697 the suite of *n*-alkanes observed in the DCM extracts of the POP gun and SPE samples at the 698 Ecurie site (Figs. 9b and 10a). Although long-chain alkanes are usually associated with plant 699 waxes, Simoneit et al. (2004) suggested that their occurrence in the chimney deposits may reflect 700 thermal processing of organic matter within the Rainbow hydrothermal environment. 701 Conversely, a direct microbial source for these longer-chain compounds cannot be precluded.

702 Combined, the UCMs and alkanes observed in the Ecurie fluids provide a very close match 703 to the organic compounds extracted from chimneys at Rainbow, suggesting a common source for 704 the organic compounds observed in the chimney deposits and the POP gun and SPE samples. 705 While most of the fluids collected for this study were sampled directly from natural orifices, at 706 Ecurie the fluids were sampled from a cavity excavated into the side of a chimney structure that 707 was weakly venting lower temperature fluids (~60 °C) at its apex (Fig. 1b). The tip of the POP 708 gun was embedded in loose chimney materials during sampling, and it is likely that some 709 chimney particles were entrained into the POP gun and SPE cartridge during sampling, and are 710 the source of the UCM and alkanes observed in the POP gun and SPE extracts. If this is the 711 case, it suggests that a substantial fraction of the UCM found in the hydrothermal chimneys 712 could be seawater-derived organic matter that has undergone alteration to remove some of the 713 lower molecular weight fraction. Whether this alteration is primarily chemical or biological 714 remains to be determined.

The source of smaller UCMs observed in POP gun samples from the higher temperature Stylo 2 and Padraig sites at Rainbow is more difficult to evaluate (Fig. 8b). The POP gun extracts for these sites also contain several PAHs, and the narrow UCM peak and lack of alkanes observed in the POP gun samples distinguishes them from the suite of organic compounds reported for chimney samples. These observations suggest that the UCM and PAH found in the extracts were present as dissolved compounds in the fluids.

Extracts of the two POP gun samples from the Guillaume site at Rainbow contained hexadecanol and a C_{18} alkenoic acid in addition to several PAH (Fig. 8c). Hexadecanol was not observed in other samples from Rainbow, but it was observed in the fluid sampled from the top of Poseidon at Lost City (Fig. 5e-g). While hexadecanol is a plausible microbial product, it is difficult to explain how this compound would occur in isolation without being accompanied by a suite of other alkanols or alkanoic acids if derived from a biological source. On the other hand, there is no obvious source of contamination that would explain the occurrence of this compound in the samples. Thus, the hexadecanol appears to be indigenous to the samples, but the ultimate source for this compound remains obscure. Similarly, the C_{18} alkenoic acid found at the Guillaume site appears to be an indigenous biological product, but its source is uncertain.

731 Overall, several PAH, particularly phenanthrene together with its methyl and dimethyl 732 homologs, were the only organic compounds consistently found in the high-temperature fluids 733 from the Rainbow field that appeared to be indigenous to the fluids (e.g., Figs. 8 and 9). 734 Although the PAHs were not quantified, they were present at very low levels that were barely 735 detectable in the extracts. Phenanthrene and other PAHs are prominent constituents of organic 736 matter that has been exposed to high-temperature alteration in hydrothermal environments 737 (Simoneit, 1988; Kawka and Simoneit, 1994; Ventura et al., 2012), suggesting that the PAH 738 observed in the fluids at Rainbow may represent organic matter that has undergone thermal 739 processing during circulation through the hydrothermal system. Because this process obscures 740 the original composition of the organic matter, it is difficult to identify the ultimate source of 741 these compounds.

742 At the Lucky Strike site, the hydrothermal fluid extracts were largely devoid of organic 743 compounds that could be attributed to the fluids with a high degree of confidence, with the 744 possible exception of trace dibenzothiophene that was observed in a couple of samples (e.g., Fig. 745 7b). The only POP gun sample obtained at Lucky Strike (J2-359-POP1) was found to contain a 746 number of fatty acid methyl esters, dominated by C_{16} and C_{18} *n*-alkanoic and *n*-alkenoic acids 747 (Fig. 8d) (note that fatty acids present in the sample were probably converted to methyl esters 748 during sample processing). These compounds were also observed in the SPE sample attached to 749 the POP gun (Fig. 10b), but not in the IGT fluid sample obtained at the same vent site (Fig. 7b). 750 The similarity of the fatty acids found in the POP gun and SPE samples to the components of 751 bacterial membranes strongly suggests that these compounds were derived from microbial 752 biomass entrained during sample collection with the POP gun rather than representing dissolved 753 organic compounds from the hydrothermal fluid.

754 In addition to an aliphatic component, deep seawater DOM has also been found to contain a 755 significant component of polycyclic aromatic compounds with attached carboxyls and other functional groups (e.g., Dittmar and Koch, 2006). It has been hypothesized that some of these 756 compounds could originate in submarine hydrothermal systems. Our results do not show 757 758 evidence for a substantial input of aromatic compounds into the deep sea from high temperature 759 hydrothermal systems, since such compounds were only detected in trace amounts. However, 760 more substantial amounts of PAH could be contributed from hydrothermal systems that are 761 buried by organic-rich sediments where fluids may discharge much higher amounts of aromatic 762 compounds (Simoneit, 1988).

763 **4.3. Comparison with previous studies**

764 The results of this study stand in sharp contrast with previous studies that have reported the 765 presence of diverse organic compounds in fluids from the Rainbow and Lost City vent fields, as 766 well as for a hydrothermal system on the East Pacific Rise (Brault et al., 1988; Konn et al., 2009, 767 2012). Konn et al. (2012) extracted small aliquots (10 ml) of hydrothermal fluids from the 768 Rainbow field using a stir-bar solid phase extraction approach, and reported detection of a 769 diverse suite of organic compounds that included alkanes, alkanoic acids, cycloalkanes, and 770 aromatic hydrocarbons. The alkanoic acids observed ranged from C_9 to C_{18} , with a predominance of even carbon number and particularly high abundances of C₁₂, C₁₄, and C₁₆ 771 acids. In contrast, we observed no alkanoic acids in the Rainbow samples, despite extraction of 772

much larger sample volumes (>10× larger volumes for the IGT fluid samples, ~100× larger for the POP gun and SPE samples). Although Konn et al. (2012) did not quantify the amount of acids in their samples, the relatively strong GC-MS response they observed for the C_{12} - C_{16} acids suggests that these compounds would have been readily detectable in our samples if they had been present at comparable levels (or even at substantially lower levels).

778 Konn et al. (2012) also reported observation of a homologous series of *n*-alkanes ranging 779 from C_9 to C_{19} , with especially high abundances of C_{10} , C_{11} , and C_{12} alkanes at Rainbow. We 780 observed *n*-alkanes in only one Rainbow sample (Ecurie), and these compounds had a much 781 higher range of carbon numbers (C_{23} - C_{33}). Alkanes in the carbon number range reported by 782 Konn et al. (2012) were not found in any of the hydrothermal fluids we sampled at Rainbow 783 (note that undecane and several methlydecanes *were* observed in all of our samples wherever 784 MeOH was used as a solvent, but these compounds were identified as contaminants from the 785 solvent and were not derived from the fluids). In agreement with our results, Konn et al. (2012) 786 also reported the observation of trace amounts of PAHs, but found a much broader spectrum of 787 compounds than the limited number observed during this study (primarily phenanthrene and 788 alkylated homologs; Table 2). Although monocyclic alkylated aromatic compounds and 789 cycloalkanes were prominent components of the suite of organic compounds reported by Konn et 790 al. (2012), the methods employed for this study were not suited for identification of compounds 791 of that type.

792 In an earlier study, Konn et al. (2009) had reported detection of a suite of compounds at Lost 793 City that were very similar to those they identified at Rainbow. Again, their results are very 794 different from those found in the present study. While Konn et al. (2009) reported alkanoic acids 795 to be present in the Lost City fluid, these compounds were only present at low levels and were 796 dominated by the C₁₂, C₁₄, and C₁₆ acids rather than the C₈, C₁₀, and C₁₂ acids found to be 797 dominant in this study. Konn et al. (2009) report carbon isotopic compositions for the C8-C12 alkanoic acids, but their reported values ($C_8 = -31.1\%$, $C_{10} = -19.6\%$, and $C_{12} = -25.5\%$) are 798 799 considerably lighter than the compositions measured for the same compounds in this study 800 (Table 3), although the values reported by Konn et al. are similar to the longer-chain alkanoic 801 acids in the Lost City chimney extract. Konn et al. (2009) also report detection of a number of 802 alkanes and PAH at Lost City, but none of those were observed in the samples analyzed for this 803 study.

804 Some key aspects of the organic compounds identified in the Rainbow and Lost City fluids 805 by Konn et al. (2009, 2012) are particularly noteworthy. First, nearly identical suites of organic compounds were reported for hydrothermal fluids in both systems, even though the fluids have 806 807 radically different chemistries and have experienced very different hydrothermal circulation 808 histories (see, for example, Fig. 9 of Konn et al., 2009). End-member hydrothermal fluids at 809 Rainbow attain temperatures well in excess of 365 °C and are among the most acidic and metal-810 rich observed in mid-ocean ridge hydrothermal systems worldwide (Table 1; Charlou et al., 811 2002). In contrast, the hydrothermal fluids discharged from the Lost City system have 812 apparently experienced maximum temperatures of about 150-250 °C during circulation through 813 the subsurface and have become strongly alkaline through the process of serpentinization (Kelley 814 et al., 2005; Proskurowski et al., 2006; Foustoukos et al., 2008). There is no obvious reason to 815 expect that the organic composition of the Rainbow and Lost City fluids should be similar, and it 816 would be truly remarkable if fluids experiencing such radically different conditions were 817 characterized by the same organic content.

818 Second, the suite of organic compounds reported by Konn et al. has some highly unusual 819 characteristics that would be difficult to explain by natural processes. For instance, the *n*-alkanes 820 are dominated by C_{10} and C_{11} compounds in both systems, with much lower amounts of other *n*-821 alkanes. Typically, natural sources of organic compounds display much broader and more 822 regular distributions of *n*-alkanes, with peaks in distribution at higher carbon number (e.g., 823 Simoneit et al., 2004; Delacour et al., 2008). Conversely, abiotic organic synthesis pathways 824 produce mixtures of compounds that display a regular decrease in alkane abundance with 825 increasing carbon number rather than having a peak at a discrete number of carbons (e.g., 826 McCollom et al., 2010). A natural source for high abundances of C_{10} and C_{11} *n*-alkanes to the 827 exclusion of lower and higher carbon number compounds is difficult to explain. Also, while homologous series of alkanes and alkanoic acids in geologic samples tend to exhibit regular 828 829 patterns of abundance with carbon number (e.g., Brault et al., 1988; Simoneit et al., 2004), the 830 homologues reported by Konn et al. (2009, 2012) exhibit no regular or consistent pattern. For 831 example, the alkanoic acids show no regular abundance with carbon number, and the C_{10} , C_{11} , and C_{12} *n*-alkanes appear to be present in roughly equal amounts while those with both fewer 832 833 (C₉) and greater (C₁₃-C₁₉) are present in substantially lower amounts (see Fig. 5 in Konn et al., 834 2012). It is somewhat difficult to conceive of a natural process that could explain the irregular 835 distributions observed for these compounds.

836 It is not clear how to reconcile the discrepancies between the results of the present study and 837 those reported for the same sites by Konn et al. (2009, 2012). The discrepancies may reflect real 838 temporal or spatial differences in the organic composition of the hydrothermal fluids that were 839 sampled, or differences in the effectiveness of the protocols to extract and detect organic 840 compounds from the fluids. However, the successful detection of organic compounds in the 841 samples from Lost City and other sites indicates that the methods used in the present study were 842 effective in extracting dissolved organic compounds when they are present, suggesting that it is 843 unlikely that our study would have missed the compounds reported by Konn et al. (2009, 2012) 844 if they had been present. Alternatively, the unusual distribution of compounds and close 845 similarities between different sites observed by Konn et al. (2009, 2012) may be an indication 846 that many of the compounds they reported are derived from an unrecognized background source 847 of contamination rather than the hydrothermal fluids. Further comparative tests of the extraction 848 and analytical protocols employed in these various studies as well as analysis of additional 849 samples of hydrothermal fluids using multiple methods are needed to resolve this issue.

850 In an earlier study, Holm and Charlou (2001) reported observation of trace amounts of C_{16} to C₂₉ *n*-alkanes in samples of high-temperature (364 °C) hydrothermal fluids from the Rainbow 851 852 field. These compounds were extracted from 50 ml aliquots of hydrothermal fluids that had been 853 stored in a cold room for a considerable length of time prior to analysis. Organic compounds 854 were extracted from the fluid with SPE units containing the same HLB sorbent used in this 855 study. In the present study, alkanes were not identified in any of the high-temperature (326-366 856 $^{\circ}$ C) hydrothermal fluids sampled at Rainbow, even in the sample from Stylo2 where more than a 857 liter of fluid was pumped through the SPE cartridge containing HLB. The only site where alkanes were observed at Rainbow was at the lower temperature Ecurie site (Figs. 9b and 10a), 858 859 but the alkanes in those samples have a substantially different carbon number distribution (C₂₂ to 860 C_{33}) from those reported by Holm and Charlou (2001), and are most likely derived from chimney particulates entrained during sampling. Consequently, we are unable to reproduce or confirm the 861 862 results of Holm and Charlou (2001).

863 The only other study to provide a detailed analysis of organic matter extracted from deep-sea hydrothermal vent fluids is Brault et al. (1988). Those authors analyzed dissolved organic 864 865 compounds in >250 °C fluid venting from a basalt-hosted hydrothermal system at 13°N on the 866 East Pacific Rise collected with a 350 ml Ti syringe sampler and extracted with a liquid-liquid method using chloroform as the solvent. The extract contained a suite of *n*-alkanes ranging in 867 868 size from C₁₇ to C₃₅ along with an unconsolidated complex mixture (UCM) of unidentified 869 compounds dominated by aliphatic components. The UCM hump in the extract was broad with a 870 peak toward the higher molecular weight end centered at about C_{29} . The alkanes exhibited a smooth distribution with a peak in abundance at about C_{28} , with a very slight preference for even 871 872 carbon number compounds. The total amount of aliphatic hydrocarbons in their hydrothermal fluid sample was $\sim 56 \ \mu g \ l^{-1}$, which was $\sim 100 \ times$ greater than levels in the local seawater. 873

The suite of organic compounds observed by Brault et al. (1988) were not observed in any of 874 875 the hydrothermal fluids examined in the present study using comparable liquid-liquid extraction 876 methods. Although Brault et al. (1988) extracted a larger volume of hydrothermal fluid than was 877 included in the IGT samples (350 ml vs. ~120 ml) analyzed here, if aliphatic hydrocarbons had 878 been present in comparable amounts to those reported by Brault et al. they should have been 879 readily detectable in the IGT extracts. Samples from the Ecurie site at Rainbow were the only 880 samples in this study where a suite of compounds similar to those reported by Brault et al. (1988) 881 were identified. These samples contained both a suite of alkanes peaking at $\sim C_{28}$ and a broad 882 UCM with a peak towards the higher molecular weight end (Figs. 9b and 10a). In the case of the 883 Ecurie sample, however, the organic compounds in the extract appear to be derived from 884 particles of chimney sulfides included in the sample rather than the hydrothermal fluid itself. 885 These observations suggest that the compounds reported by Brault et al. (1988) may also have 886 been derived from chimney particulates entrained into the sampler during fluid collection, which 887 common occurrence for the large syringe samplers used in that study.

888 4.4. Abiotic organic compounds

889 One of the primary reasons that deep-sea hydrothermal systems have become strong 890 candidates as the site of the origin of life is that conditions within these systems appear to be 891 highly favorable for the abiotic synthesis of organic compounds (e.g., Shock, 1990, 1992; Shock 892 and Schulte, 1998; McCollom and Seewald, 2007). Ultramafic-hosted hydrothermal systems are 893 viewed as particularly favorable environments for abiotic synthesis owing to the strongly 894 reducing conditions and highly elevated concentrations of H_2 that develop during 895 serpentinization of ultramafic rocks (Berndt et al., 1996; McCollom and Seewald, 2001, 2013). Indeed, a substantial body of evidence has accumulated that the methane and other light 896 897 hydrocarbons found in fluids discharging from the ultramafic-hosted deep-sea hydrothermal 898 vents at Lost City and Rainbow have a predominantly abiotic origin (Prokurowski et al., 2008; 899 Charlou et al., 2002, 2010). Formate has also been identified as a product of abiotic carbon 900 reduction at Lost City (Lang et al., 2010).

Whether other, more-complex organic compounds with an abiotic origin can be identified in deep-sea hydrothermal fluids remains an open issue. Holm and Charlou (2001) reported detection of C_{16} - C_{29} *n*-alkanes in hydrothermal fluids from the Rainbow system that they attributed to abiotic formation through Fischer-Tropsch-type synthesis in the hydrothermal system. Konn et al. (2009) also speculated that some of the *n*-alkanes and *n*-alkanoic acids they observed in extracts of Lost City and Rainbow hydrothermal fluids might have an abiotic origin.

In the present study, no evidence was found for the presence of abiotic organic compounds in any of the hydrothermal fluid extracts. Although linear alkanes are the primary products of

909 Fischer-Tropsch-type abiotic organic synthesis (e.g., McCollom and Seewald, 2007; McCollom 910 et al., 2010), alkanes were observed at only one site during this study (Ecurie at Rainbow) and 911 those compounds were likely derived from particulate matter rather than the fluid. Furthermore, 912 while Fischer-Tropsch products characteristically exhibit a log-linear decrease in abundance with 913 increasing carbon number (e.g., McCollom and Seewald, 2006; McCollom et al. 2010), the 914 alkanes observed at the Ecurie site show a bell-shaped distribution centered at ~C₂₈ with no compounds below C_{23} (Figs. 9b and 10a). Thus, the distribution of these hydrocarbons does not 915 916 resemble the distribution that would be expected for abiotically formed compounds. As 917 suggested by Simoneit et al. (2004), thermal alteration of organic precursors within the 918 hydrothermal system is the most likely source for these compounds. Fischer-Tropsch synthesis 919 also produces alkanoic acids, but the abundance of those compounds would also be expected to 920 follow a regular log-linear decrease with carbon number with no even-odd carbon number 921 preference (e.g., McCollom et al., 1999; McCollom et al. 2010). While alkanoic acids were 922 found in the Lost City fluids and chimney extracts, the strong even-carbon-number preference 923 for these compounds indicates they are biological products.

924 The absence of extractable abiotic hydrocarbons in the hydrothermal fluids should not 925 necessarily be completely unexpected. Assuming that abiotic hydrocarbons in deep-sea 926 hydrothermal fluids would follow a distribution similar to typical Fischer-Tropsch products, 927 extrapolation from the concentrations of light hydrocarbons observed in the Lost City and Rainbow systems (which are in the nmol kg^{-1} range) to predict the expected concentrations of the 928 extractable higher hydrocarbons ($\geq C_{10}$) suggests that these compounds should be present at 929 levels of 1 pmol kg⁻¹ or less (Fig. 12). At such low concentrations, it would be very difficult to 930 931 detect abiotic hydrocarbons even if they could be concentrated significantly during extraction by 932 SPE or other methods. Other abiotic high-molecular-weight organic compounds (alkanoic acids, 933 etc.) might be expected to have similarly low concentrations, suggesting that it may be extremely 934 challenging to detect the presence of these compounds in deep-sea hydrothermal fluids.

935 **5. CONCLUDING REMARKS**

936 The results of this study illustrate many of the difficulties faced in trying to detect solvent 937 extractable hydrocarbons and other higher molecular weight dissolved organic compounds in 938 end-member deep-sea hydrothermal fluids. These compounds, if they are present at all, appear 939 to occur at dissolved concentrations that are at or below the detection limit of commonly used 940 extraction and analysis methods. Additional detection of organic compounds may require use of 941 methods to concentrate the compounds from large amounts of fluid prior to analysis. 942 Furthermore, because of the very low levels at which these compounds occur, it is essential to 943 fully evaluate potential background sources of organic compounds that could easily be mistaken 944 for components dissolved in the fluids, including those from sources such as chimney 945 particulates, biomass, and contaminants introduced during sample handling. The POP gun and 946 SPE sorbent approaches used in this study show some promise for concentrating organic 947 compounds and lowering detection limits, but clearly the level of background compounds from 948 the sorbent phases needs to be significantly reduced to maximize the utility of this method. If 949 this can be accomplished, the ability to concentrate organic compounds from several liters or 950 more of hydrothermal fluids may allow even trace compounds to be detected and characterized, 951 and may ultimately reveal additional details of the sources and fate of organic matter in deep-sea 952 hydrothermal environments.

953

954 Acknowledgements – This research was supported by the NSF Ocean Sciences directorate 955 through grants MGG-OCE 0550800 to T.M.M. and MGG-OCE 0549829 to J.S.S. and C.R.G. 956 Additional support for completion of analyses and manuscript preparation was provided by the 957 Alfred P. Sloan Foundation though the Deep Carbon Observatory. The authors thank the captain 958 and crew of the R/V Roger Revelle and the crew of the Jason ROV who made collection of the 959 samples possible. The authors also thank chief scientist Anna-Louise Reysenbach and the other 960 members of the KNOX18RR expedition for their help and cooperation during the cruise. The 961 authors are grateful for helpful comments from three anonymous reviewers and AE Jeff Alt that 962 greatly improved the manuscript. 963

964

965	REFERENCES
966	
967	Barnes I., Rapp J. B., O'Neill J. R., Sheppard R. A., and Gude A. J., III (1972) Metamorphic
968	assemblages and the direction of flow of metamorphic fluids in four instances of
969	serpentinization. <i>Contrib. Mineral. Petrol.</i> 35 , 263-276.
970	Berndt M. E., Allen D. E. and Seyfried W. E., Jr. (1996) Reduction of CO ₂ during
971	serpentinization of olivine at 300°C and 500 bar. <i>Geology</i> 24 , 351-354.
972	Bradley A. S. and Summons R. E. (2009) Multiple origins of methane at the Lost City
973	Hydrothermal Field. <i>Earth Planet. Sci. Lett.</i> 297 , 34-41.
974	Bradley A. S., Hayes J. M., and Summons R. E. (2009) Extraordinary ¹³ C enrichment of diether
975	lipids at the Lost City Hydrothermal Field indicates a carbon-limited ecosystem. Geochim.
976	<i>Cosmochim. Acta</i> 73 , 102-118.
977	Brault M., Simoneit B. R. T., Marty J. C. and Saliot A. (1988) Hydrocarbons in waters and
978	particulate material from hydrothermal environments at the East Pacific Rise, 13°N. Org.
979	<i>Geochem.</i> 12 , 209-219.
980	Brazelton W. J., Schrenk M. O., Kelley D. S., and Baross J. A. (2006) Methane- and sulfur-
981	metabolizing microbial communities dominate the Lost City hydrothermal field ecosystem.
982	Appl. Environ. Microbiol. 72, 6257-6270.
983	Charlou J. L., Donval J. P., Jean-Baptiste P., Dapoigny A. and Rona P. A. (1996) Gases and
984	helium isotopes in high temperature solutions sampled before and after ODP Leg 158
985	drilling at TAG hydrothermal field (26°N MAR). Geophys. Res. Lett. 23, 3491-3494.
986	Charlou J. L., Donval J. P., Douville E., Jean-Baptiste P., Radford-Knoery J., Fouquet Y.,
987	Dapoigny A., and Stievenard M. (2000) Compared geochemical signatures and the
988	evolution of Menez Gwen (37°50'N) and Lucky Strike (37°17'N) hydrothermal fluids,
989	south of the Azores Triple Junction on the Mid-Atlantic Ridge. Chem. Geol. 171, 49-75.
990	Charlou J. L., Donval J. P., Fouquet Y., Jean-Baptiste P. and Holm N. (2002) Geochemistry of
991	high H ₂ and CH ₄ vent fluids issuing from ultramafic rocks at the Rainbow hydrothermal
992	field (36° 14'N, MAR). Chem. Geol. 191, 345-359.
993	Charlou J. L., Donval J. P., Konn C., Ondréas H. and Fouquet Y. (2010) High production and
994	fluxes of H ₂ and CH ₄ and evidence of abiotic hydrocarbon synthesis by serpentinization in
995	ultramafic-hosted hydrothermal systems on the Mid-Atlantic Ridge. In Diversity of
996	Hydrothermal Systems on Slow-spreading Ocean Ridges (ed. P. Rona, C. Devey, J. Dyment,
997	B. Murton) Geophysical Monograph Series. pp. 265–296.
998	Delacour A., Früh-Green G. L., Bernasconi S. M., Schaeffer P. and Kelley D. S. (2008) Carbon
999	geochemistry of serpentinites in the Lost City Hydrothermal System (30°N, MAR).
1000	Geochim. Cosmochim. Acta 72 , 3681-3702.
1001	Deming J. W., and Baross J. A. (1993) Deep-sea smokers: Windows to a subsurface biosphere?
1002	Geochim. Cosmochim. Acta 57 , 3219–3230.
1003	Dittmar T. and Koch B. P. (2006) Thermogenic organic matter dissolved in the abyssal ocean.
1004 1005	Mar. Chem. 102 , 208-217.
1005	Dittmar T. and Paeng J. (2009) A heat-induced molecular signature in marine dissolved organic matter. <i>Nat. Geosci.</i> 2 , 175-179.
1000	Feth J. H., Rogers S. M., and Roberson C. E. (1961) Aqua de Ney, California, a spring of unique
1007	chemical character. <i>Geochim. Cosmochim. Acta</i> 22, 75-86.
1008	Foustoukos D. I. and Seyfried W. E., Jr. (2004) Hydrocarbons in hydrothermal vent fluids:
1010	The role of chromium-bearing catalysts. <i>Science</i> 304 , 1002-1005

- Foustoukos D. I., Savov I. P., and Janecky D. R. (2008) Chemical and isotopic constraints on
 water/rock interactions at the Lost City hydrothermal field, 30°N Mid-Atlantic Ridge. *Geochim. Cosmochim. Acta* 72, 5457-5474.
- Foustoukos D. I., Pester N. J., Ding K., and Seyfried W. E., Jr. (2009) Dissolved carbon species
 in associated diffuse and focused flow hydrothermal vents at the Main Endeavour Field,
 Juan de Fuca Ridge: Phase equilibria and kinetic constraints. *Geochem. Geophys. Geosyst.*
- 1017 **10**, Q10003, doi:10.1029/2009GC002472.
- Fuchida S., Mizuno Y., Masuda H., Toki T., and Mikita H. (2014) Concentrations and
 distributions of amino acids in black and white smoker fluids at temperatures over 200 °C. *Org. Geochem.* 66, 98-106.
- German C. R. and Seyfried W. E., Jr. (2014) Hydrothermal Processes. In *Treatise on Geochemistry*, 2nd Edition (ed. H. D. Holland and K. K.Turekian). Vol. 8, pp. 191-233.
- Haberstroh P. R. and Karl D. M. (1989) Dissolved free amino acids in hydrothermal vent
 habitats of the Guaymas Basin. *Geochim. Cosmochim. Acta* 53, 2937-2945.
- Hertkorn N., Harir M., Koch B. P., Michalke B., and Schmitt-Kopplin P. (2013) High-field NMR
 spectroscopy and FTICR mass spectrometry: powerful discovery tools for the molecular
 level characterization of marine dissolved organic matter. *Biogosci.* 10, 1583-1624.
- Holm N. G., and Charlou J. L. (2001) Initial indications of abiotic formation of hydrocarbons in
 the Rainbow ultramafic hydrothermal system, Mid-Atlantic Ridge. *Earth Planet. Sci. Lett.* **191**, 1-8.
- Horiuchi T., Takano Y., Ishibashi J., Marumo K., Urabe T., and Kobayashi K. (2004) Amino
 acids in water samples from deep sea hydrothermal vents at Suiyo Seamount, Izu-Bonin
 Arc, Pacific Ocean. Org. Geochem. 35, 1121-1128.
- Huang Y., Wang Y., Alexandre M. R., Lee T., Rose-Petruck C., Fuller M., and Pizzarello S.
 (2005) Molecular and compound-specific isotopic characterization of monocarboxylic acids in carbonaceous chondrites. *Geochim. Cosmochim. Acta* 69, 1073-1084.
- Kawka O. E., and Simoneit B. R. T. (1994) Hydrothermal pyrolysis of organic matter in
 Guaymas Basin: 1. Comparison of hydrocarbon distributions in subsurface sediments and
 seabed petroleums. *Org. Geochem.* 22, 947-978.
- Kelley D. S., Karston J. A., Früh-Green G. L., Yoerger D. R., Shank T. M., Butterfield D. A.,
 Hayes J. M., Schrenk M. O., Olson E. J., Proskurowski G, Jakuba M., Bradley A., Larson
 B., Ludwig K., Glickson D., Buckman K., Bradley A. S., Brazelton W. J., Roe K., Elend M.
- J., Delacour A., Bernasconi S. M., Lilley M. D., Baross J. A., Summons R. E., Sylva S. P.
 (2005) A serpentinite-hosted ecosystem: The Lost City hydrothermal field. *Science* 307, 1428-1434.
- Klevenz V., Sumoodur A., Ostertag-Henning C., and Koschinsky A. (2010) Concentrations and
 distributions of dissolved amino acids in fluids from Mid-Atlantic Ridge hydrothermal vents.
 Geochem. J. 44, 387-397.
- Konn C., Charlou J. L., Donval J. P., Holm N. G., Dehairs F. and Bouillon S. (2009)
 Hydrocarbons and oxidized organic compounds in hydrothermal fluids from Rainbow and
 Lost City ultramafic-hosted vents. *Chem. Geol.* 258, 299-314.
- 1052 Konn C., Charlou J. L., Donval J. P., and Holm N. G. (2012) Characterisation of dissolved
- 1053 organic compounds in hydrothermal fluids by stir bar sorptive extraction-gas
- 1054 chromatography-mass spectrometry. Case study: the Rainbow field (36°N, Mid-Atlantic
- 1055 Ridge). Geochem. Transact. 13:8.

- Lang S. Q., Butterfield D. A., Schulte M., Kelley D. S., and Lilley M. D. (2010) Elevated
 concentrations of formate, acetate and dissolved organic carbon found at the Lost City
 hydrothermal field. *Geochim. Cosmochim. Acta* 74, 941-952.
- Lang, S. Q., Früh-Green G. L., Bernasconi S. M., Lilley M. D., Proskurowski G., Méhay S., and
 Butterfield D. A. (2012) Microbial utilization of abiogenic carbon and hydrogen in a
 serpentinite-hosted system. *Geochim. Cosmochim. Acta* 92, 82-99.
- Lang S. Q., Früh-Green G. L., Bernasconi S. M., and Butterfield D. A. (2013) Sources of
 organic nitrogen at the serpentinite-hosted Lost City hydrothermal field. *Geobiology* 11,
 154-169.
- Legras J. L., Erny C., Le Jeune C., Lollier M., Adolphe Y., Demuyter C., Delobel P., Blondin B.,
 and Karst F. (2010) Activation of two different resistance mechanisms in *Saccharomyces cerevisiae* upon exposure to octanoic and decanoic acids. *Appl. Environ. Microbiol.* 76,
 7526-7535.
- Lilley, M. D., J. A. Baross, and L. I. Gordon (1983) Reduced gases and bacteria in hydrothermal
 fluids: The Galapagos Spreading Center and 21°N East Pacific Rise. In *Hydrothermal Processes at Seafloor Spreading Centers* (ed. P. A. Rona et al.) Plenum Press. pp. 411–447.
- Lilley M. D., Butterfield D. A., Olson E. J., Lupton J. E., Macko S. A., and McDuff R. E. (1993)
 Anomalous CH₄ and NH₄⁺ concentrations at an unsedimented mid-ocean-ridge
 hydrothermal system. *Nature* J2-364, 45–47.
- 1075 Martens C. S. (1990) Generation of short chain organic acid anions in hydrothermally altered 1076 sediments of the Guaymas Basin, Gulf of California. *Appl. Geochem.* **5**, 71–76.
- McCollom T. M. and Seewald J. S. (2001) A reassessment of the potential for reduction of
 dissolved CO₂ to hydrocarbons during serpentinization of olivine. *Geochim. Cosmochim.* Acta 65, 3769-3778.
- McCollom T. M. and Seewald J. S. (2006) Carbon isotope composition of organic compounds
 produced by abiotic synthesis under hydrothermal conditions. *Earth Planet. Sci. Lett.* 243,
 74-84.
- McCollom, T.M., and Seewald J. S. (2007) Abiotic synthesis of organic compounds in deep-sea
 hydrothermal environments. *Chem. Rev.*, **107**, 382-401.
- McCollom T. M. and Seewald J. S. (2013) Serpentinites, hydrogen, and life. *Elements*, 9, 129134.
- McCollom T. M., Ritter G., and Simoneit B. R. T. (1999) Lipid synthesis under hydrothermal
 conditions by Fischer-Tropsch-type reactions. *Origins Life Evol. Biosphere* 29, 153-166.
- McCollom T. M., Sherwood Lollar B[•], Lacrampe-Couloume G. and Seewald, J. S. (2010) The
 influence of carbon source on abiotic organic synthesis and carbon isotope fractionation
 under hydrothermal conditions. *Geochim. Cosmochim. Acta* 74, 2717-2740.
- Méhay S., Früh-Green G. L., Lang S. Q., Bernasconi S. M., Brazelton W. J., Schrenck M. O.,
 Schaeffer P. and Adam P. (2013) Record of archaeal activity at the serpentinite-hosted Lost
 City Hydrothermal Field. *Geobiology* 11, 570-592.
- Pasini V., Brunelli D., Dumas P., Sandt C., Frederick J., Benzerara, Bernard S., and Ménez B.
 (2013) Low temperature hydrothermal oil and associated biological precursors in
 serpentinites from Mid-Ocean Ridge. *Lithos* 178, 84-95.
- Proskurowksi G., Lilley M. D., Kelley D. S., and Olson E. J. (2006) Low temperature volatile
 production at the Lost City Hydrothermal Field, evidence from a hydrogen stable isotope
 geothermometer. *Chem. Geol.* 229, 331-343.

- Proskurowksi G., Lilley M. D., Seewald J. S., Früh-Green G. L., Olson E. J., Lupton J. E.,
 Sylva S. P., and Kelley D. S. (2008) Abiogenic hydrocarbon production at Lost City
 Hydrothermal Field. *Science* 319, 604-607
- Reeves E. P., McDermott J. M., and Seewald J. S. (2014) The origin of methanethiol in midocean ridge hydrothermal fluids. *Proc. Nat. Acad, Sci. USA* 111, 5474-5479.
- Schrenk M. O., Kelley D. S., Bolton S. A., and Baross J. A. (2006) Low archaeal diversity
 linked to subseafloor geochemical processes at the Lost City hydrothermal filed, MidAtlantic Ridge. *Environ. Microbiol.* 6, 1086-1095.
- Seewald J. S., Doherty K. W., Hammar T. R., and Liberatore S. P. (2002) A new gas-tight
 isobaric sampler for hydrothermal fluids. *Deep-Sea Res. I* 49, 189-196.
- Seyfried W. E., Jr., Foustoukos D. I. and Fu. Q. (2007) Redox evolution and mass transfer during
 serpentinization: An experimental and theoretical study at 200 °C, 500 bar with implications
 for ultramafic-hosted hydrothermal systems at Mid-Ocean Ridges. *Geochim. Cosmochim. Acta* 71, 3872-3886.
- Seyfried W. E., Jr., Foustoukos D. I. and Fu. Q. (2011) Vent fluid chemistry of the Rainbow
 hydrothermal system (36°N, MAR): Phase equilibria and in situ pH controls on subseafloor
 alteration processes. *Geochim. Cosmochim. Acta* **75**, 1574-1593.
- Shock E. L. (1990) Geochemical constraints on the origin of organic compounds in hydrothermal
 systems. *Origins Life Evol. Biospheres* 20, 331-367.
- Shock E. L. (1992) Chemical environments of submarine hydrothermal systems, *Origins Life Evol. Biospheres* 22, 67-107.
- Shock E. L., and Schulte M. D. (1998) Organic synthesis during fluid mixing in hydrothermal
 systems. J. Geophys. Res. 103, 28,513-28,517.
- Simoneit B. R. T. (1988) Petroleum generation in submarine hydrothermal systems: An update.
 Can. Mineral. 26, 827-840.
- Simoneit B. R. T., Lein A.Yu., Peresypkin V. I., and Osipov G. A. (2004) Composition and
 origin of hydrothermal petroleum and associated lipids in the sulfide deposits of the
 Rainbow Field (Mid-Atlantic Ridge at 36°N). *Geochim. Cosmochim. Acta* 68, 2275-2294.
- Ventura G. T., Simoneit B. R. T., Nelson R. K., and Reddy C. M. (2012) The composition, origin
 and fate of complex mixtures in the maltene fractions of hydrothermal petroleum assessed
 by comprehensive two-dimensional gas chromatography. Org. Geochem. 45, 48-65.
- 1132 Von Damm K. L., Bray A. M., Buttermore L. G., and Oosting S. E. (1998) The geochemical
 1133 controls on vent fluids from the Lucky Strike vent field, Mid-Atlantic Ridge. *Earth Planet*.
 1134 Sci. Lett. 160, 521-536.
- Welhan J. A. and Craig H. (1983) Methane, hydrogen, and helium in hydrothermal fluids at
 21°N on the East Pacific Rise. In *Hydrothermal Processes at Seafloor Spreading Centers* (
 ed. P. A. Rona, et al.), pp. 391–409. Plenum Press.
- Welhan J. A. and Lupton J. E. (1987) Light-hydrocarbon gases in Guaymas Basin hydrothermal
 fluids—Thermogenic versus abiogenic origin. *Bull. Am. Assoc. Petrol. Geol.*, **71**, 215–223.
- Yuen G. U. and Kvenvolden K. A. (1973) Monocarboxylic acids in Murray and Murchison
 carbonaceous meteorites. *Science* 246, 301-303.
- 1142
- 1143

Site/Vent	Temp.	Lat.	U	Sample ID			pН	Mg	H_2	ΣCO_2	CH ₄
	(°C)	(N)		IGT	POP gun	SPE(phase)	(25°C)				
	Rainbo	ow (ultramafic	-hosted)								
Guillaume	359	36° 13.752'	33° 54.222'	J2-352-IGT7	J2-352-POP1, J2-352-POP2		3.24	1.9	15.8	21.2	2.03
Regner	326	36° 13.751'	33° 54.232'	J2-353-IGT7			3.45	8.0	13.5	21.1	1.86
Padraig	350	36° 13.766'	33° 54.173'	J2-353-IGT6	J2-353-POP1, J2-353-POP2		3.43	5.4	16.1	19.6	1.79
Stylo 1	366	36° 13.781'	33° 54.070'	J2-354-IGT7			3.28	2.1	15.0	24.6	2.10
Ecurie	191	36° 13.748'	33° 54.188'	J2-354-IGT6	J2-354-POP1, J2-354-POP2	J2-354-SPE(HLB)	2.99	7.4	9.4	18.7	1.79
Stylo 2	359	36° 13.777'	33° 54.072'	J2-355-IGT7	J2-355-POP1, J2-355-POP2	J2-355-SPE(C8), J2-355-SPE(HLB)	3.27	2.6	16.5	20.9	2.10
	Lucky	Strike (basalt-	hosted)								
2608(top) US4 Crystal Medea	323 299 306 251	37° 17.508' 37° 17.276' 37° 17.450' 37° 17.659'	32° 16.675' 32° 16.529' 32° 16.920' 32° 16.900'	J2-356-IGT6 J2-357-IGT6 J2-358-IGT6 J2-359-IGT6	J2-359-POP1 [#]	J2-359-SPE(HLB)#	3.86 4.75 3.64 3.68	0.9 29.5 1.5 2.2	0.048 0.053 0.041 0.063	34.7 133 117 98	1.01 0.74 0.82 0.89
		ity (ultramafic		020071010	02 007 1 01 1	<u>, , , , , , , , , , , , , , , , , , , </u>	0100		01000	20	0.05
Beehive	91	30° 7.431'	42° 7.203'	J2-360-IGT6			10.6	0.1	10.4	0.18	1.9
Beehive	90	30° 7.434'	42° 7.202'	J2-361-IGT6	J2-361-POP1, J2-361-POP2	J2-361-SPE(HLB), J2-361-SPE(C8)	10.6	1.0	10.4	0.18	1.9
Top of Poseidon	51	30° 7.430'	42° 7.203'	J2-362-IGT6		、 <i>/</i>	10.8	2.6	-	-	-
Seawater (Lucky Seawater (Lost Ci Seawater (TAG)				J2-364-IGT6	J2-356-POP1 J2-362-POP1	J2-356-SPE(HLB) J2-362-SPE(C8)	- - -	- - -	- - -	- - -	- - -

Table 1. Summary of samples collected for analysis of organic compounds, with selected chemical information.

Dissolved concentrations of Mg, H₂, Σ CO₂, and CH₄ in mmol kg⁻¹. Σ CO₂ is total dissolved CO₂ (=CO_{2(aq)} + HCO₃⁻ + CO₃²⁻). "-" = not measured. [#]For this sample, fluid was processed through the POP gun for only 20 minutes; for all other POP gun samples, fluids were pumped for about 60 minutes.

	Rainbow						Lucky Strike				Lost City	
Site: Sample types:	Guillaume (I,P,S)	Regner (I)	Padraig (I,P,S)	Stylo1 (I)	Ecurie (I,P,S)	Stylo2 (I,P,S)	2608 (I)	US4 (I)	Crystal (I)	Medea (I,P,S)	Beehive (I,P,S)	Poseidon (I)
Carboxylic acids												
Octanoic acid	-	-	-	-	-	-	-	-	-	-	Р	Р
Decaonoic acid	-	-	-	-	-	-	-	-	-	-	Р	Р
Dodecanoic acid	-	-	-	-	-	-	-	-	-	-	Р	Р
Tetradecanoic acid	-	-	-	-	-	-	-	-	-	В	-	-
Hexadecanoic acid	-	-	-	-	-	-	-	-	-	В	-	-
Hexadecenoic acid	-	-	-	-	-	-	-	-	-	В	-	-
Octadecanoic acid	-	-	-	-	-	-	-	-	-	В	-	-
Octadecenoic acid	Р	-	-	-	-	-	-	-	-	В	-	-
Doeicosanoic acid	-	-	-	-	-	-	-	-	-	В	-	-
Aliphatic hydrocarbons												
<i>n</i> -Alkanes [range]	-	-	-	-	CS	-	-	-	-	-	-	-
					$[C_{23}-C_{33}]$							
Squalene	-	-	-	-	-	-	-	-	-	-	-	Р
Aromatic hydrocarbons												
Phenanthrene	Р	-	Р	-	Р	Р	-	-	-	-	-	-
Methylphenanthrenes	Р	-	Р	-	Р	Р	-	-	-	-	-	-
C ₂ -phenanthrenes	-	-	Р	-	Р	Р	-	-	-	-	-	-
Methylfluorene	Р	-	-	-	-	-	-	-	-	-	-	-
Dibenzothiophene	-	-	-	-	-	-	-	-	Р	Р	-	-
Other compounds												
Cholesterol	-	-	-	-	-	-	-	-	-	-	-	Р
Hexadecanol	Р	-	-	-	-	-	-	-	-	-	-	Р
Trithiolane	-	-	-	-	-	-	-	-	-	-	-	Р
UCM	-	-	-	-	CS	Р	-	-	-	-	-	-

Table 2. Summary of organic compounds detected in hydrothermal vent fluid samples.

Compounds that were positively identified and inferred to be indigenous to the fluid are designated with a "P", while compounds that were detected but interpreted to come from other sources are identified designated with a "CS" (chimney sulfide source) or "B" (biomass source). "-" = compound not detected in fluid sample. Sample types: I = IGT sampler, P = POP gun, S = SPE cartridge.

Sample	J2-360- IGT6	J2-361- IGT6	J2-361- SPE(HLB)					
	μmol acid/L							
C_8	0.72	0.53	0.77					
C_{10}	0.47	0.37	0.54					
C ₁₂	0.28	0.23	0.31					
		- μmol C/L						
C_8	5.7	4.2	6.2					
C_{10}	4.7	3.7	5.4					
C ₁₂	3.4	2.7	3.7					
Total	13.8	10.6	15.3					
		$\delta^{13}C$ (‰)						
C_8	-5.9	-7.6						
C_{10}	-4.3	-2.8						
C ₁₂	-4.8	-3.6						

Table 3. Abundances[†] and isotopiccomposition of n-alkanoic acids in Lost Cityhydrothermal fluids samples.

[†]Concentrations are given for the acids and for the equivalent number of carbon atoms.

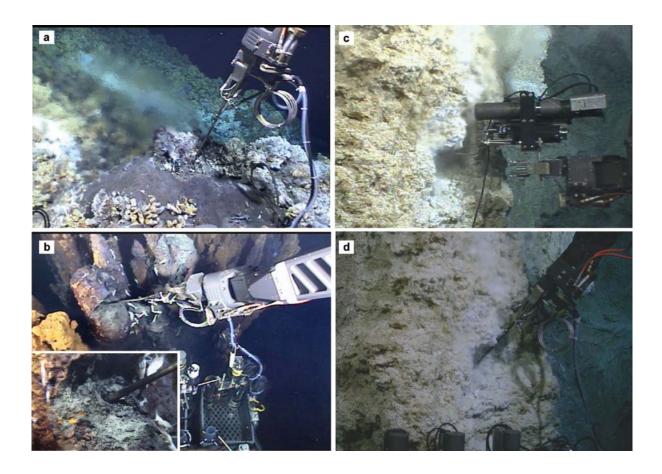


Figure 1. Representative images of fluid sampling at the deep-sea hydrothermal vent sites. (a) Sampling with the POP gun at the Medea site at Lucky Strike. (b) Sampling with the POP gun at the Ecurie site at Rainbow. The fluid at this site was obtained from a cavity excavated into the side of an older, less active chimney structure. Inset shows closeup of tip of POP gun in side of chimney. (c, d) Fluid sampling at the Beehive site at Lost City with the isobaric gas-tight sampler (c) and POP gun (d).

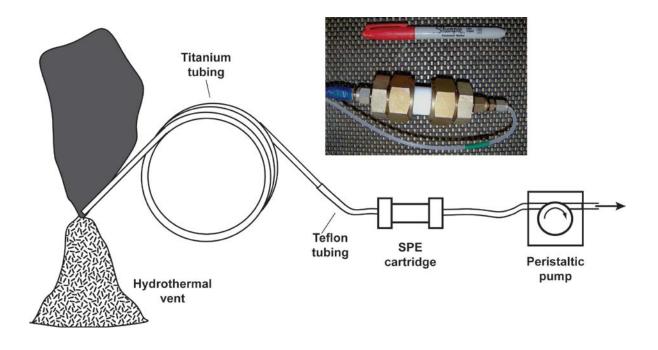


Figure 2. Schematic drawing of the peristaltic organic pump fluid sampling device (POP gun). The sampler consists of a coil of titanium tubing (4 m long) connected through Teflon tubing to a cartridge filled with solid phase extraction (SPE) microbeads, which is then connected to a peristaltic pump with silicon tubing. During operation, the tip of the sampler is placed in a hydrothermal vent, and the peristaltic pump draws fluid through the tubing and SPE cartridge. The inset shows an image of the SPE cartridge, consisting of SPE microbeads within a 2.54 cm diameter Teflon tube capped with porous Teflon frits and sealed at the ends with stainless steel Swagelock fittings.

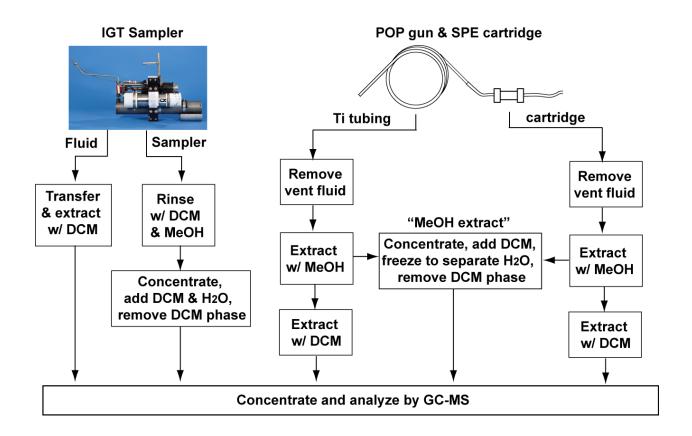


Figure 3. Flow diagram summarizing methods used to extract and prepare organic compounds for analysis from the different sampling methods used in this study.

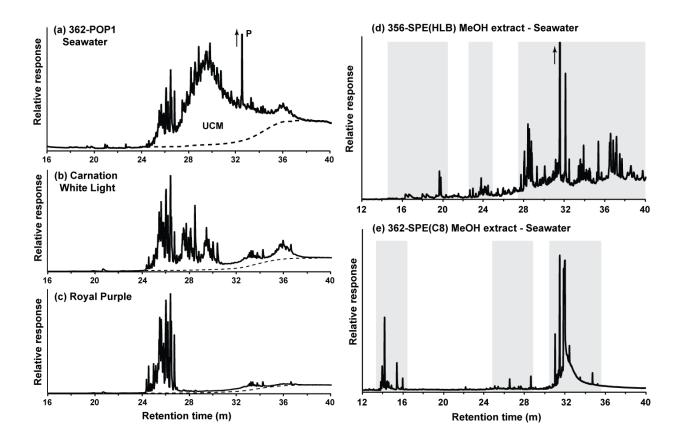


Figure 4. Total ion chromatograms from GC-MS analysis of potential sources of background organic compounds that could have contaminated the hydrothermal fluid samples. (a) MeOH extract of seawater sample from the vicinity of the TAG hydrothermal field obtained with the POP gun, showing a large unresolved complex mixture of compounds (UCM). Results for a seawater sample taken in the vicinity of the Lucky Strike hydrothermal vent field (J2-356-POP1) were very similar. (b,c) Oils used for lubrication and pressure compensation on ROV *Jason*. These oils also display a UCM, but they are distinctly different from the seawater UCM. (d,e) MeOH extracts of deep seawater samples processed using the HLB and C8 SPE cartridges. Gray areas highlight prominent peaks for organic compounds that were present in all samples processed through these cartridges, including Fisher H₂O blanks, and are inferred to be contaminants extracted from the SPE microbeads. In all but a few cases, the mass fragmentation patterns for these compounds did not match any compounds in the reference database and could not be identified.

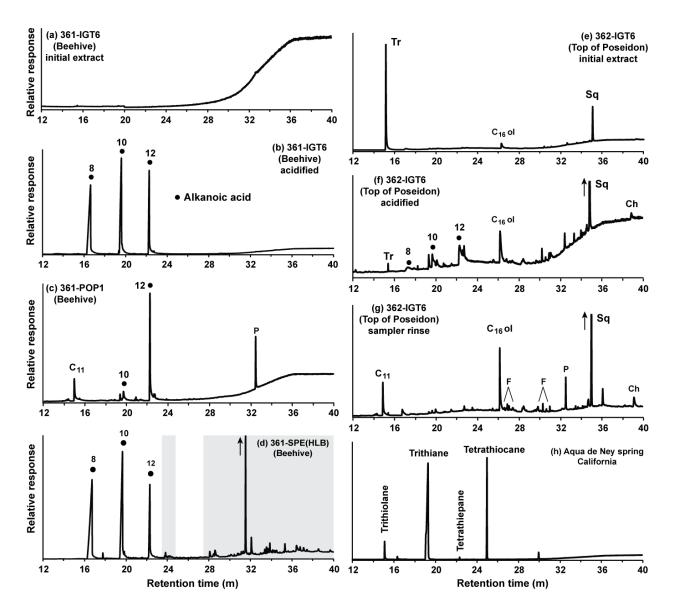


Figure 5. Total ion chromatograms for samples from the Lost City hydrothermal system. (a,e) Initial DCM extracts for samples from Beehive and Top of Poseidon (ToP). The Beehive sample shows no peaks above the instrumental baseline, while the ToP sample has peaks for trithiolane (Tr), hexadecanol (C_{16} ol), and squalene (Sq). (b,f) Second DCM extracts from Beehive and ToP samples following acidification of the fluid. The Beehive sample exhibits prominent peaks for octanoic, decanoic and dodecanoic acids (C_8 , C_{10} and C_{12} , respectively), which are also evident in smaller amounts in the ToP sample along with cholesterol (Ch). Sample J2-360-IGT6 produced results identical to the Beehive sample shown. (c) Analysis of POP gun MeOH extract from Beehive site. The other POP gun sample for this site produced the same result. (d) Sample from Beehive vent processed with the HLB SPE cartridge and extracted with MeOH. Peaks outlined by gray areas are inferred to be contaminants because they were present in all MeOH extracts of the HLB cartridges (see Fig. 4d and Supplemental Fig. S1). (g) Solvent rinse of IGT sampler from ToP site. (h) Analysis of DCM extract of highly alkaline fluid (pH = 11.8) discharged from Aqua de Ney spring, northern California. The spring fluid includes trithiolane and several other cyclic C-S compounds (trithiane, tetrathiepane, tetrathiocane). The sample was

collected in a glass bottle and extracted using the same procedures as the IGT samples. Undecane (C_{11}) comes from the methanol solvent, compounds labeled "F" are from the Fluorolube lubricant, and "P" is dioctyl phthalate. Unlabeled peaks could not be identified.

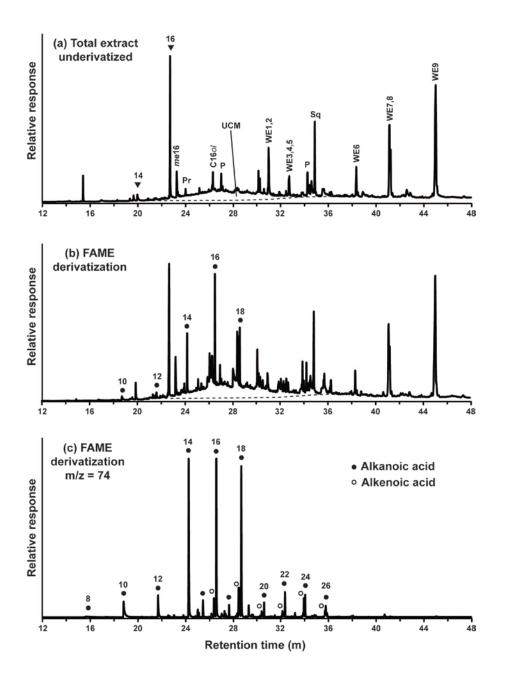


Figure 6. Chromatograms of Lost City chimney extracts. (a) Total ion chromatogram for analysis of total extract. Dashed line shows instrument baseline, and area marked UCM is an unresolved complex mixture of compounds. Peaks labeled "WE" are wax esters, and other labeled peaks are *n*-alkanes (inverted triangles, labeled with carbon number), methylhexadecane (*me*16), pristane (Pr), hexadecaol (C16*ol*), and squalene (Sq). (b) Total ion chromatogram of total extract derivatized to convert fatty acids to methyl esters (FAME). Only peaks for methyl esters of the C₁₀-C₁₈ alkanoic acids are labeled; other peaks correspond to those labeled in (a) and (c). (c) Mass fragmentogram for mass-to-charge ratio (m/z) of 74, characteristic for FAME.

Labels identify methyl esters of *n*-alkanoic and *n*-alkenoic acids, with number of carbon atoms specified for even-numbered compounds. Alkenoic acids elute just before alkanoic acid of same carbon number (positions of double bonds were not determined). "P" = phthalates. Other abbreviations: WE1,2 = C_{10} - C_{12} ester + C_8 - C_{14} ester; WE3,4,5 = C_{12} - C_{12} ester + C_{10} - C_{14} ester + C_8 - C_{14} ester; WE6 = C_{14} - C_{16} ester; WE7,8 = C_{14} - C_{18} ester + C_{16} - C_{16} ester; WE9 = C_{16} - C_{18} ester.

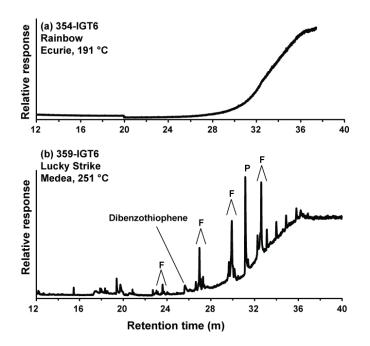


Figure 7. Total ion chromatograms for selected GC-MS analysis of extracts of IGT fluid samples. Families of peaks marked with an "F" represent contamination from the Fluorolube lubricant used in the samplers, and the peak labeled "P" is a phthalate. (a) Extract of fluid from the Ecurie site at Rainbow, showing an absence of peaks above the instrumental baseline. Most other fluid extracts were similar. (b) Extract of fluid from the Medea site at Lucky Strike. Unlabeled peaks in the chromatograms did not match any compounds in the spectral library, and are presumably additional contaminants introduced during sample processing.

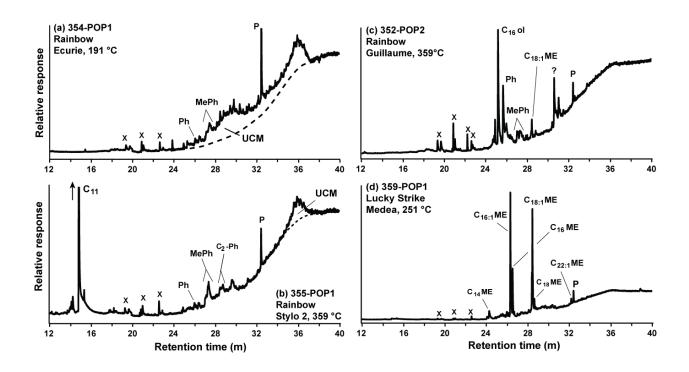


Figure 8. Total ion chromatograms for representative MeOH extracts of POP gun samples. The dashed lines in the chromatograms represent the instrument baseline. (a) Sample from the Ecurie vent at Rainbow (191 °C; J2-354-POP1) showing a small, broad UCM extending from ~25 to 38 min and peaks for phenanthrene (Ph) and methylphenanthrenes (MePh). The sample also contains several C₂-phenanthrenes (C₂-Ph), but peaks for these compounds are obscured by the UCM. The other POP gun sample at this site (J2-354-POP2) produced very similar results. (b) Sample from the Stylo 2 site at Rainbow (359 °C; J2-355-POP1) showing small peaks for several PAH including Ph, MePh, and C₂-Ph, as well as a small UCM centered at about 36 min. Extracts for other POP gun samples at the high-temperature Stylo 2 and Padraig sites at Rainbow showed similar trace levels of aromatic compounds. (c) Sample from the 359 °C Guillaume site at Rainbow. In addition to Ph and MePh, peaks corresponding to hexadecanol (C160l) and octadecenoic acid methyl ester (C18:1ME) are present. The other POP gun sample at this site produced identical results. (d) Sample from the Medea vent at Lucky Strike (251 °C), including methyl esters of a number of alkanoic and alkenoic acids (C_rME and $C_{r1}ME$, respectively, where x is the carbon number). Undecane (C_{11}) comes from the methanol solvent. Peaks marked with an "X" were unidentifiable compounds present in all samples and are presumably contaminants introduced during sample processing (see Supplemental Fig. S1). "P" = dioctyl phthalate.

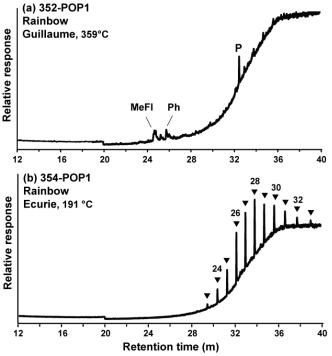


Figure 9. Total ion chromatograms for selected DCM extracts of POP gun samples. (a) Sample from the 359 °C Guillaume site at Rainbow, showing trace amounts of methylfluorene (MeFl) and phenanthrene (Ph). (b) Sample from the Ecurie site at Rainbow, showing peaks for a homologous series of *n*-alkanes with 23-33 carbon atoms (inverted triangles; labels show carbon chain length). Peak labeled "P" is dioctylphthalate. DCM extracts for other POP gun samples showed no organic compounds attributable to the hydrothermal fluids.

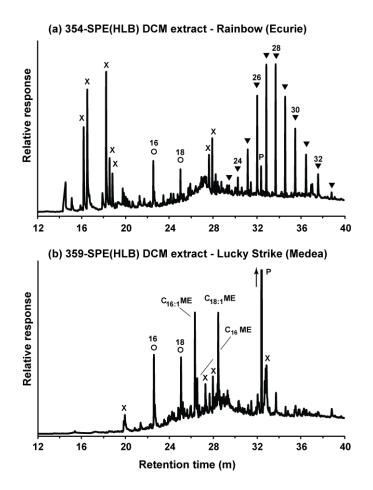


Figure 10. Total ion chromatograms for selected HLB SPE cartridge samples extracted with DCM. (a) Sample from the Ecurie site at Rainbow, showing peaks for a homologous series of *n*-alkanes with 23-33 carbon atoms (inverted triangles; numbers show carbon chain length). (c) Sample from Medea vent at Lucky Strike. Included in the extract are methyl esters of several alkanoic and alkenoic acids (C_x ME and $C_{x:1}$ ME, respectively, where *x* is the carbon number). Prominent peaks labeled with "X" as well as the broad hump of compounds centered at ~27.5 m are present in all DCM extracts of the HLB cartridges and are inferred to be contaminants from the SPE microbeads (see Supplemental Fig. S1 for examples). Note that the sample shown in (b) was inadvertently allowed to evaporate to dryness, resulting in loss of compounds at <19 m. Also identified as ubiquitous contaminants from the SPE microbeads are C₁₆ and C₁₈ *n*-alkenes (labeled with open circles). Peak labeled "P" is dioctylphthalte.

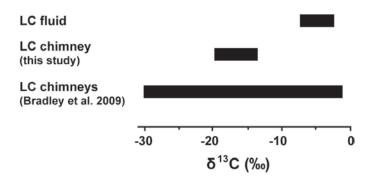


Figure 11. Range of carbon isotopic compositions for *n*-alkanoic acids measured for fluid and chimney samples from Lost City in this study, and reported for multiple active chimney samples from Lost City by Bradley et al. (2009).

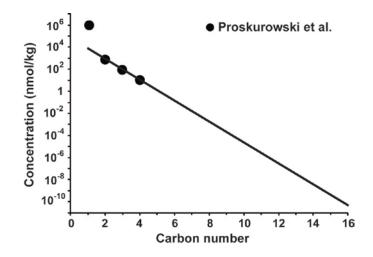


Figure 12. Predicted concentrations of linear alkanes as function of carbon number (line) based on extrapolation of measurements of C_2 - C_4 alkanes measured at the Beehive vent at Lost City by Proskurowski et al. (2008) (circles). Line calculated assuming a log-linear decrease in alkane abundance with increasing carbon number typical of Fischer-Tropsch-type organic synthesis reactions.