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1 **Real time imaging, forecasting and management of**
2 **human-induced seismicity at Preston New Road,**
3 **Lancashire, England**

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15 **ABSTRACT**

16 *Earthquakes induced by subsurface fluid injection pose a significant issue across a range of industries.*
17 *Debate continues as to the most effective methods to mitigate the resulting seismic hazard.*
18 *Observations of induced seismicity indicate that the rate of seismicity scales with the injection volume,*
19 *and that events follow the Gutenberg-Richter distribution. These two inferences permit us to populate*
20 *statistical models of the seismicity, and extrapolate them to make forecasts of the expected event*
21 *magnitudes as injection continues. Here we describe a shale gas site where this approach was used in*
22 *real time to make operational decisions during hydraulic fracturing operations.*
23 *Microseismic observations revealed the intersection between hydraulic fracturing and a pre-existing*
24 *fault or fracture network that became seismically active. While “red light” events, requiring a pause to*
25 *the injection program, occurred on several occasions, the observed event magnitudes fell within*
26 *expected levels based on the extrapolated statistical models, and the levels of seismicity remained*
27 *within acceptable limits as defined by the regulator. To date, induced seismicity has typically been*
28 *regulated using retroactive Traffic Light Schemes. This study shows that the use of high quality*
29 *microseismic observations to populate statistical models that forecast expected event magnitudes can*
30 *provide a more effective approach.*

31

32

33 1. Introduction

34 Human-induced seismicity is becoming an increasingly controversial topic. It is well known
35 that activities such as mining and water impoundment can lead to felt seismicity, but
36 increasingly activities such as geothermal energy (Grigoli *et al.*, 2018), underground storage
37 of waste such as CO₂ or water (Keranen *et al.*, 2014), production from conventional
38 hydrocarbon reservoirs (e.g. Segall, 1989) and hydraulic stimulation of shale gas reservoirs
39 (Bao and Eaton, 2016), are attracting concern from the public, regulators and operators.

40 The stimulation of fractures by injecting water at high-pressure is a technique used to create
41 conductive fracture networks in low-permeability reservoir rocks. Hydraulic fracture
42 stimulation is widely used in the commercial production of hydrocarbons, and also to develop
43 engineered geothermal systems. Use of this method has become more prominent in the past
44 decade, associated primarily with the shale gas boom (Wang and Krupnick, 2013) in North
45 America.

46 If hydraulic fractures intersect a pre-existing fault that is near to its critical stress state, the
47 increase in pore pressure can reduce the effective normal stress, declamping the fault and
48 creating induced seismicity. Such cases are relatively rare: Atkinson *et al.* (2016) estimate
49 that only 0.3% of wells in British Columbia and Alberta, a region with some of the highest
50 levels of hydraulic fracturing-induced seismicity (HF-IS), are associated with induced events
51 larger than magnitude 3. Nonetheless, the issue of induced seismicity is a concern for the
52 petroleum and geothermal industries, and will likely be of concern to other nascent industries,
53 such as carbon capture and storage, as well (e.g., Verdon, 2014).

54 Debate continues with regards to the most effective methods to mitigate HF-IS, and what
55 regulations should be applied. To date, regulators have typically imposed Traffic Light
56 Schemes (TLSs) whereby the operator reduces, pauses or stops injection if the magnitude of
57 the largest event exceeds a specified threshold. TLS thresholds have varied significantly in
58 different jurisdictions (Bosman *et al.*, 2006; Baisch *et al.*, 2019): for example, in Alberta the
59 red light is set at $M = 4$, whereas in the United Kingdom (U.K.) the red light is set at $M = 0.5$,
60 a difference in earthquake moment of over 175,000 times.

61 The simple TLSs currently used by hydraulic fracturing regulators are essentially retroactive
62 in nature, because the operator takes actions after an event has occurred. In some case studies,
63 seismicity has been observed to continue, and increase in magnitude, after injection has
64 ceased (e.g., Häring *et al.*, 2008; Clarke *et al.*, 2014). These post-injection increased-
65 magnitude events, known as “trailing events”, pose an issue for TLSs because they compel

66 the regulator to set thresholds that may be substantially lower than the actual magnitude they
67 wish to avoid. Hence operations may be stopped even though levels of seismicity are well
68 below that which might be considered hazardous.

69 It is therefore desirable to manage and mitigate induced seismicity in real time, as operations
70 proceed. For example, injection volumes or pressures could be reduced (e.g., Kwiatek et al.,
71 2019), or stimulation can be directed away from areas showing fault reactivation. Here we
72 show a successful example of managing HF-IS with a recently acquired dataset from a shale
73 gas operation in the UK.

74

75 **1.1 Using microseismic data for decision-making to mitigate induced seismicity**

76 The TLSs described by Bosman et al. (2016) and Baisch et al. (2019) that are currently used
77 to regulate hydraulic fracturing stipulate decisions based solely on the magnitude of the
78 largest events. This is a rational option if monitoring is provided by national or regional
79 seismometer networks, where monitoring stations may be 10s of km from the site (e.g.,
80 Clarke *et al.*, 2014; Friberg *et al.*, 2014; Skoumal *et al.*, 2015; Schultz *et al.*, 2015). In such
81 cases only the larger events may be detected, and hypocentral locations and focal mechanisms
82 may be poorly constrained. Hence the only reliable, well-constrained data are the magnitudes
83 of the larger events.

84 However, it is not uncommon for operators to deploy microseismic monitoring, where
85 downhole geophone arrays (Maxwell *et al.*, 2010) or dense surface arrays (Chambers *et al.*,
86 2010) are able to detect very low magnitude “microseismic” events. High-quality
87 microseismic monitoring may record thousands or even hundreds of thousands of events with
88 very precise locations, spanning several orders of magnitude, provided in real time during
89 operations (e.g., Zinno et al., 1998). These data will be highly relevant for understanding the
90 risks posed by HF-IS. However, such data is not utilized by the relatively simple TLSs
91 currently being applied by hydraulic fracturing regulators (Bosman et al., 2016).

92 There are two primary ways by which microseismic observations can be used to guide
93 decisions to mitigate induced seismicity. Firstly, microseismic data can be used to detect and
94 characterise the interactions between hydraulic fractures and pre-existing faults (Maxwell *et*
95 *al.*, 2008; Maxwell *et al.*, 2009; Wessels *et al.*, 2011; Kettlety *et al.*, 2019; Igonin *et al.*, 2019;
96 Eyre et al., 2019). Microseismic events during hydraulic fracturing typically occur in clusters
97 extending from the well perpendicular to the minimum horizontal stress, tracking the growth
98 of the hydraulic fractures and mapping the extent of the stimulated reservoir volume. If a fault

99 is intersected events may begin to line up along the structure, allowing it to be identified and
100 mapped (e.g., Maxwell *et al.*, 2008; Wessels *et al.*, 2011; Hammack *et al.*, 2014; Kettlety *et*
101 *al.*, 2019 Igonin *et al.*, 2019; Eyre *et al.*, 2019). In many cases fault reactivation can also be
102 identified by a decrease in Gutenberg and Richter (1944) *b* values (e.g., Maxwell *et al.*, 2009;
103 Verdon and Budge, 2018; Kettlety *et al.*, 2019), or by an increase in the rate of
104 microseismicity relative to the injection rate (e.g., Maxwell *et al.*, 2008; Verdon and Budge,
105 2018).

106 If a fault is identified during injection, then an operator can re-design their injection program
107 to avoid further interacting with the fault. This can be achieved, for example: by skipping
108 stages along a horizontal well; by changing the planned injection rates or volumes; or by
109 altering the properties of the injected fluid (for example a more viscous fluid will carry more
110 proppant while travelling less distance into the formation). Alternatively, Hofmann *et al.*
111 (2018) have proposed adopting a “cyclic soft stimulation” program, where repeated injection
112 is conducted at significantly lower rates. Zang *et al.* (2019) have demonstrated this approach
113 for experimental-scale injection tests. However, the results from application to an industrial-
114 scale project (Hofmann *et al.*, 2019) are more ambiguous, as the Pohang geothermal project,
115 South Korea, at which this method was applied, went on to experience one of the largest
116 injection-induced events ever recorded (Grigoli *et al.*, 2018). Moreover, for shale gas
117 hydraulic fracturing applications, it is not clear that such a low-rate injection program would
118 result in effective proppant placement into a shale formation.

119 Microseismic data can also be used to make forecasts of the expected event magnitudes
120 during stimulation. Induced seismicity has been observed to follow the Gutenberg and Richter
121 (G-R hereafter) distribution (van der Elst *et al.*, 2016), with the total number of events
122 (Shapiro *et al.*, 2010; Mignan *et al.*, 2017) or the cumulative seismic moment released (Hallo
123 *et al.*, 2014) being scaled to the cumulative injection volume. As such, expected event
124 magnitudes can be forecast by characterising these relationships for the site in question, and
125 then extrapolating them to the planned injection volume. This approach has shown significant
126 promise when applied in a pseudo-prospective manner (e.g., Verdon and Budge, 2018).

127 These concepts have produced more advanced approaches to mitigate induced seismicity. For
128 example, Mignan *et al.*, (2017) propose an adaptive Traffic Light Scheme (ATLS), whereby
129 the daily rate of seismicity is scaled to the injection rate (as per Shapiro *et al.* (2010)), with
130 the addition of a post-injection relaxation time that describes trailing effects. Event
131 magnitudes are then determined from a G-R distribution, from which risk-based decisions can
132 be made. Broccardo *et al.* (2017) extended the Mignan *et al.* (2017) approach by providing a

133 Bayesian framework within which the key parameters can be estimated. However, to our
134 knowledge this approach has not yet been applied in real time to an active project.

135 Kwiatek *et al.* (2019) present an example of such methods being applied in real time to a deep
136 geothermal project near Helsinki, Finland. They found that the observed seismicity scaled
137 with injection parameters, allowing them to adjust the injection program to ensure that the
138 levels of seismicity remained within the limits imposed by the regulator. The success of the
139 type of approach demonstrated by Kwiatek *et al.* (2019), and the continued refinement of
140 proposed adaptive TLSs (e.g., Mignan *et al.*, 2017; Broccardo *et al.*, 2017), provides the
141 opportunity to move beyond the simple TLSs currently in common usage. However, their
142 effectiveness must be demonstrated extensively in real time scenarios such that regulators
143 gain confidence in their application.

144

145 **1.2 A case study from northwest England**

146 In this paper we report on the Preston New Road PNR-1z well, Lancashire, U.K., operated by
147 Cuadrilla Resources Ltd (CRL hereafter). This was the first U.K. onshore well to be
148 hydraulically fractured since a government review of HF-IS seismicity was concluded in
149 2012. As such it is the subject of regular national media attention (e.g., Webster, 2018) and
150 debate in the national parliament (Hansard, 2018). Given the high levels of public scrutiny,
151 the site was extensively monitored both by CRL, and by independently-funded organisations
152 such as the British Geological Survey (BGS). This monitoring included groundwater, surface
153 water, air quality, and traffic movements, as well as the induced seismicity monitoring
154 described here. Extensive baseline surveys were conducted for all of the above, so that any
155 change from the pre-operational conditions could be identified.

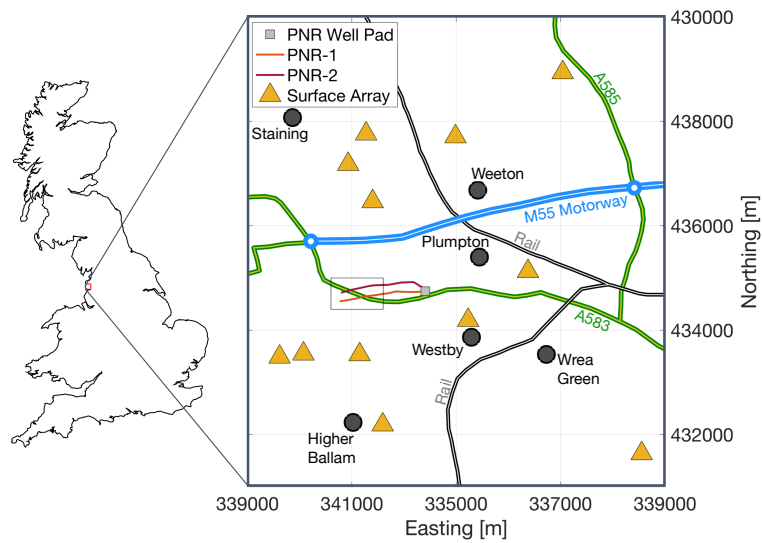
156 Given public concerns about HF-IS in the U.K., CRL took proactive measures to mitigate
157 induced seismicity, guided by microseismic observations as outlined above. Here we provide
158 a brief description of the operations conducted at the site, then show how microseismic data
159 were used to identify and map the interaction between hydraulic fractures and a fault, and to
160 forecast expected event magnitudes as the injection progressed. This information allowed
161 CRL to adjust their injection program, ensuring that levels of seismicity did not exceed the
162 overall objectives set by the regulator, as well as providing an increased understanding of
163 more proactive measures that could be applied in future as alternatives to simplistic TLSs.

164

165 **2. Description of the Preston New Road Site**

166 The Preston New Road PNR-1z well targets the Carboniferous Lower Bowland Shale at a
167 depth of approximately 2,300 m. The lateral portion of the well extends 780 m in a westward
168 direction (Figure 1). A sliding-sleeve completion was used, with 41 individual sections
169 spaced at 17.5 m intervals. CRL planned to stimulate each of these sleeves with 400 m³ of
170 slickwater, placing 50 tonnes of proppant per sleeve. Stimulation was carried out in two
171 periods (Figure 2), firstly from 15th October to 2nd November, and then from 8th to 17th
172 December 2018.

173

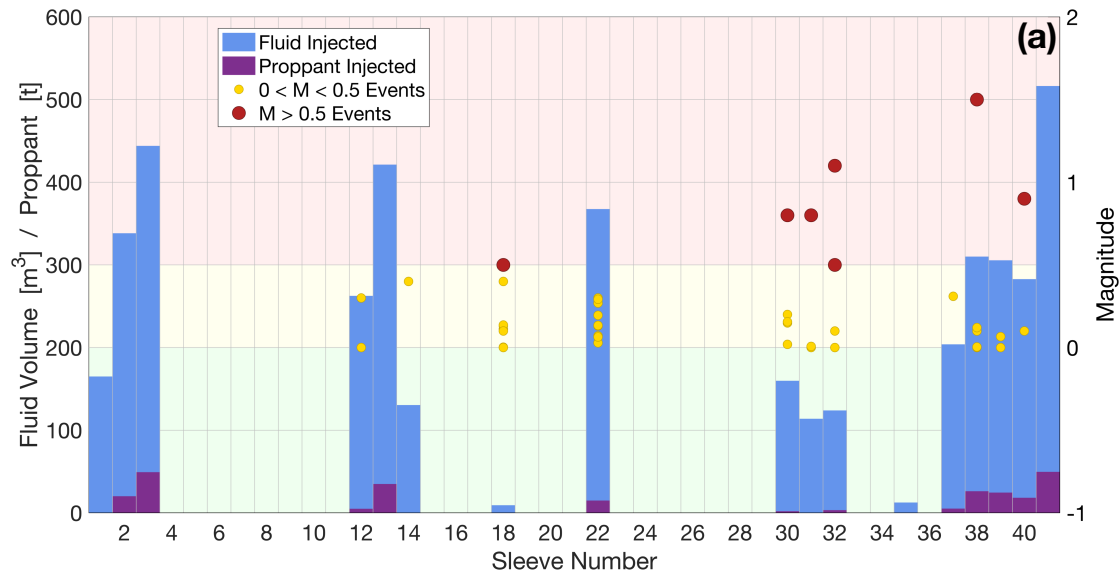


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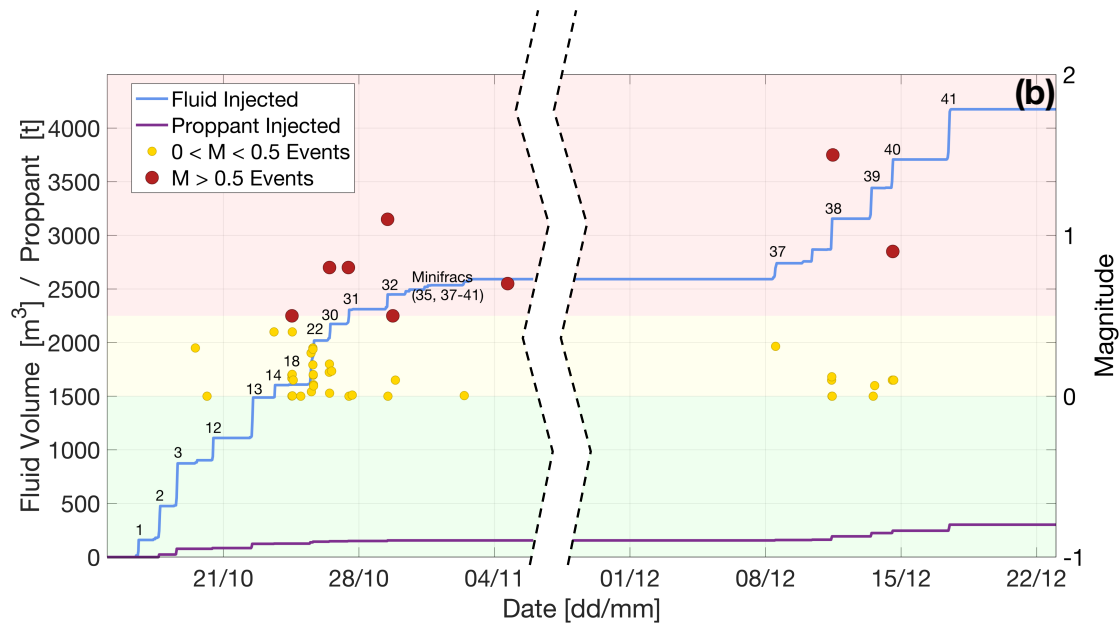
175 *Figure 1: Map of operations at Preston New Road showing the positions of the drilling pad*
176 *and horizontal tracks of PNR-1z and PNR-2, and the positions of the surface monitoring*
177 *stations. The black box marks the area of interest shown in subsequent figures. Major roads,*
178 *rail links, and nearby villages are also marked. Coordinates are U.K. Grid Reference.*

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182

183 *Figure 2: Overview of injection into PNR-1z. (a) shows the volume of fluid (blue) and mass of*
 184 *proppant (purple) injected into each sleeve. It also shows all $M > 0$ TLS events (yellow and*
 185 *red dots) that occurred during or after injection into each sleeve. (b) shows cumulative fluid*
 186 *volume (blue) and proppant mass (purple) injected as a function of time, again showing the*
 187 *occurrence of TLS events. The numbering in (b) shows the sleeve being injected. The*
 188 *background colours show the TLS green, amber and red magnitude thresholds.*

189

190 2.1 U.K. Regulations for Induced Seismicity

191 In the U.K., HF-IS is regulated by the Oil and Gas Authority (OGA). The OGA's objective is
 192 to minimize the number of events felt at the surface by the public, and to avoid the possibility
 193 of events capable of causing damage to nearby buildings or infrastructure (Oil and Gas

194 Authority, 2018). U.K. standards for ground vibrations from other activities such as quarry
195 blasting, construction equipment and industrial machinery are provided by British Standard
196 BS 7385-2. This sets a peak ground velocity (PGV) threshold, above which may cause
197 cosmetic damage such as cracking of plaster, of 15 mm/s (at lower frequencies such as would
198 be expected from induced seismicity). Using ground motion prediction equations (Akkar *et*
199 *al.*, 2014), for hypocentral depths equivalent to expected depths of hydraulic fracturing and
200 making conservative assumptions for ground conditions, this threshold is approximately
201 equivalent to a magnitude of $M = 3.5$. Therefore, the OGA's objective could be reasonably
202 translated as minimizing the number of events that have magnitudes $2 < M < 3$, and avoiding
203 events that have magnitudes $M > 3.5$.

204 To regulate HF-IS the OGA currently applies a TLS with a red-light threshold of $M = 0.5$
205 (Green *et al.*, 2012), for which the operator must stop injection, reduce the pressure in the
206 well, perform well integrity checks, and wait at least 18 hours before resuming injection. This
207 is by some margin the most stringent level for ground motion applied to any industrial
208 activity that we are aware of. The $M = 0.5$ red-light threshold is 175 times smaller than the M
209 $= 2$ events that the scheme seeks to minimize, and over 30,000 times smaller than the $M > 3.5$
210 events that the scheme seeks to avoid. This disparity exists to mitigate the risk posed by
211 trailing events, where event magnitudes may continue to increase after injection has been
212 stopped (see Mignan *et al.* (2017) for an attempt to forecast trailing event populations). This
213 TLS was applied to stimulation of the PNR-1z well, and the restrictive nature of this scheme
214 had a significant impact on the operations: only 17 of the planned 41 stages were injected,
215 and of these only 2 injected the 50 tonnes of proppant that was planned. However, only 2
216 events were reported by the British Geological Survey (BGS) as being felt, and ground
217 motions remained well below the levels at which damage might be expected. Therefore,
218 overall the operation complied with the regulator's objective to minimise felt seismicity and
219 avoid damaging seismicity.

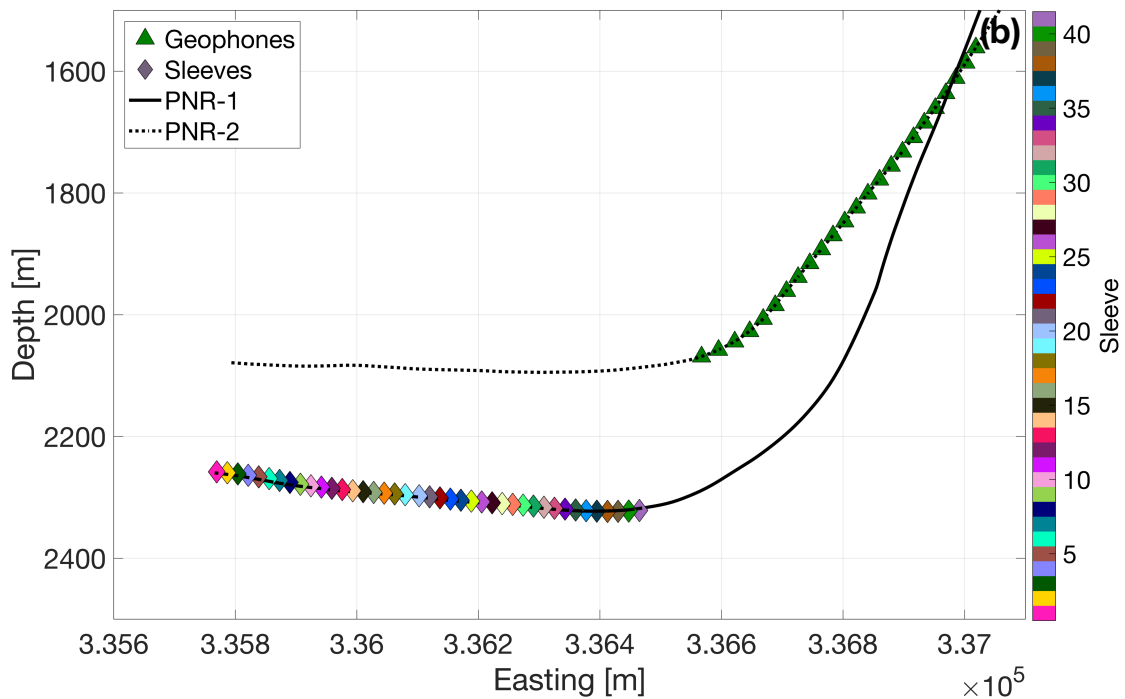
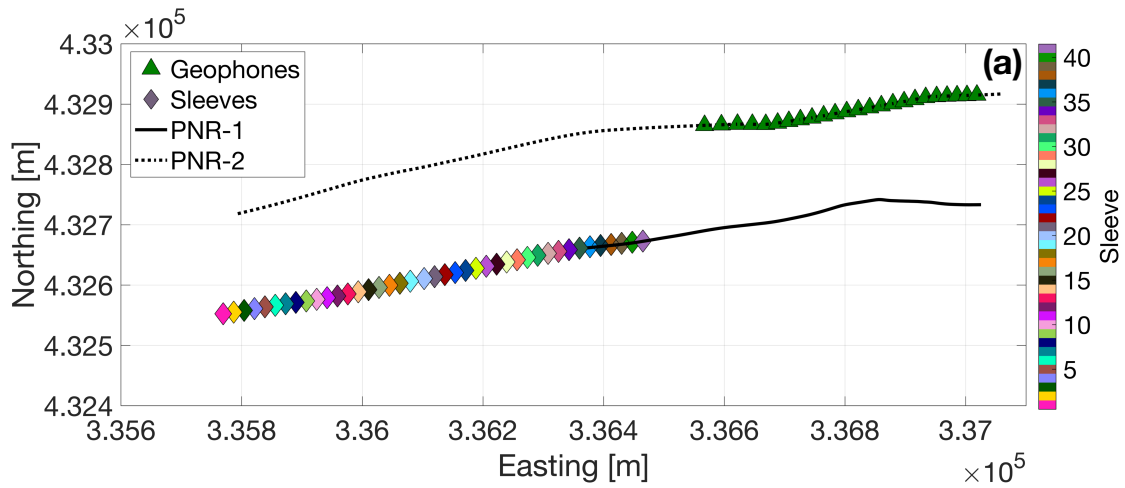
220

221 **2.2 Real time seismic and microseismic monitoring**

222 Two systems were used in combination to monitor induced seismicity at Preston New Road.
223 Both of these systems provided event locations and magnitudes in real time (typically within
224 1-4 minutes of event occurrence) computed by a processing contractor (Schlumberger). To
225 administer the TLS an array of 8 sensors including 2 broadband seismometers and 6
226 geophones (4.5 Hz instruments) was deployed at the surface, augmented by 4 broadband
227 seismometers deployed by the BGS (Figure 1). During real time monitoring the surface array
228 identified 54 events with a minimum magnitude of $M_L = -0.8$. The surface array provided

229 sufficient coverage such that focal mechanisms could be determined for 9 of the largest
230 events during real time monitoring.

231 Microseismicity was recorded using an array of 24 geophones (15 Hz instruments) placed in
232 the build section (where the well deviates from vertical to horizontal) of the adjacent PNR-2
233 well, 200 m shallower and 220 m northeast of the nearest sleeve in PNR-1z (Figure 3). This
234 array reported over 39,000 events in real time, with a minimum magnitude of $M_w = -3.0$.



237 *Figure 3: Map (a) and cross-section (b) of the downhole monitoring array deployed in well*
238 *PNR-2, and the sleeves through which injection was conducted in PNR-1z.*

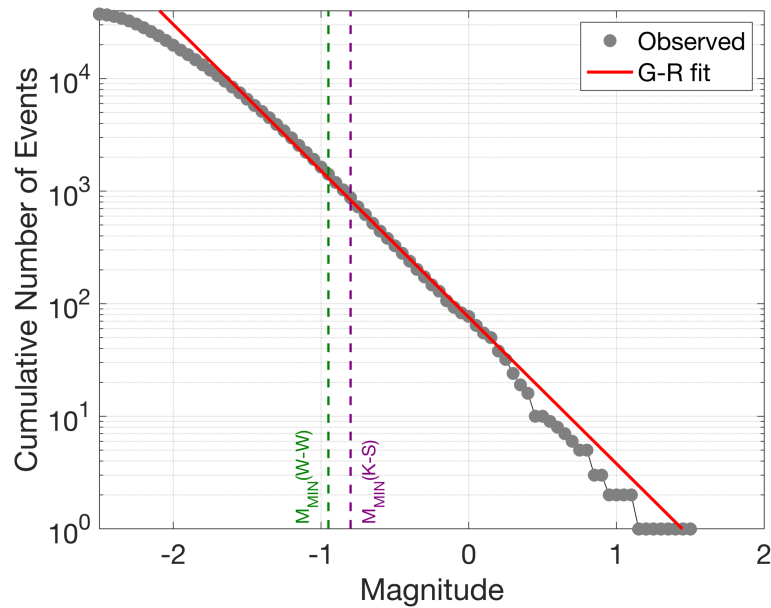
239

240 2.3 A note on magnitudes

241 Measurements of magnitudes for small events can be challenging (Kendall *et al.*, 2019). Two
242 different magnitude scales were in use during real time operations at Preston New Road. The
243 U.K. TLS regulations mandate the use of a local magnitude scale with a correction applied to
244 account for the small source-receiver distances (Butcher *et al.*, 2017; Luckett *et al.*, 2019).
245 Therefore, magnitudes from the surface array were reported as M_L values. However, these M_L
246 scales are calibrated using surface stations, implicitly including free-surface effects and near-
247 surface attenuation, so this M_L scale is not calibrated for downhole instruments. Instead the
248 downhole events were reported as M_W values. While a direct comparison and conversion
249 between the two scales might seem like an obvious solution (e.g. Edwards and Douglas,
250 2014), in practice this was more challenging. The surface array recorded the largest 54 events,
251 so only these events had reported M_L values. However, many of these larger events produced
252 subsurface motions that were beyond the dynamic range of the downhole instruments, and so
253 accurate downhole M_W values could not be determined for these events. Hence, there is only a
254 small subset of events which are large enough such that a robust M_W value can be computed
255 using the surface array, but not too large such that a robust M_W value can also be computed
256 using the downhole stations, thereby enabling a comparison to be made.

257 Work is ongoing to resolve the observed M_L and M_W values. However, the need for rapid
258 decision-making meant that this information was not used during real time operations.
259 Instead, we used M_L values for the 54 events that were reported by the surface array, and M_W
260 values for the remaining events. Clearly this solution was far from optimal. However, we note
261 that doing so does not produce anomalies or unusual behaviour if the overall magnitude-
262 frequency distribution is examined (Figure 4), suggesting that this approach was reasonable in
263 this case. However, in future cases this issue should be addressed by ensuring that moment
264 magnitudes are reported by both array types, and that relationships to convert between
265 downhole M_W values and surface M_L values are calibrated. In Figure 4 we fit a G-R
266 distribution to entire event catalogue using the Aki (1965) maximum likelihood approach,
267 computing the magnitude of completeness, M_{MIN} , using both the Wiemer and Wyss (2000)
268 formulation with an acceptance threshold of 95%, which gave $M_{MIN} = -0.95$, and by using a
269 Kolmogorov-Smirnov test with a 10% significance threshold (e.g., Clauset *et al.*, 2009;
270 Williams and Le Calvez, 2013), which gave $M_{MIN} = -0.8$. In both cases, the resulting G-R
271 parameters were $a = 1.9$ and $b = 1.3$.

272



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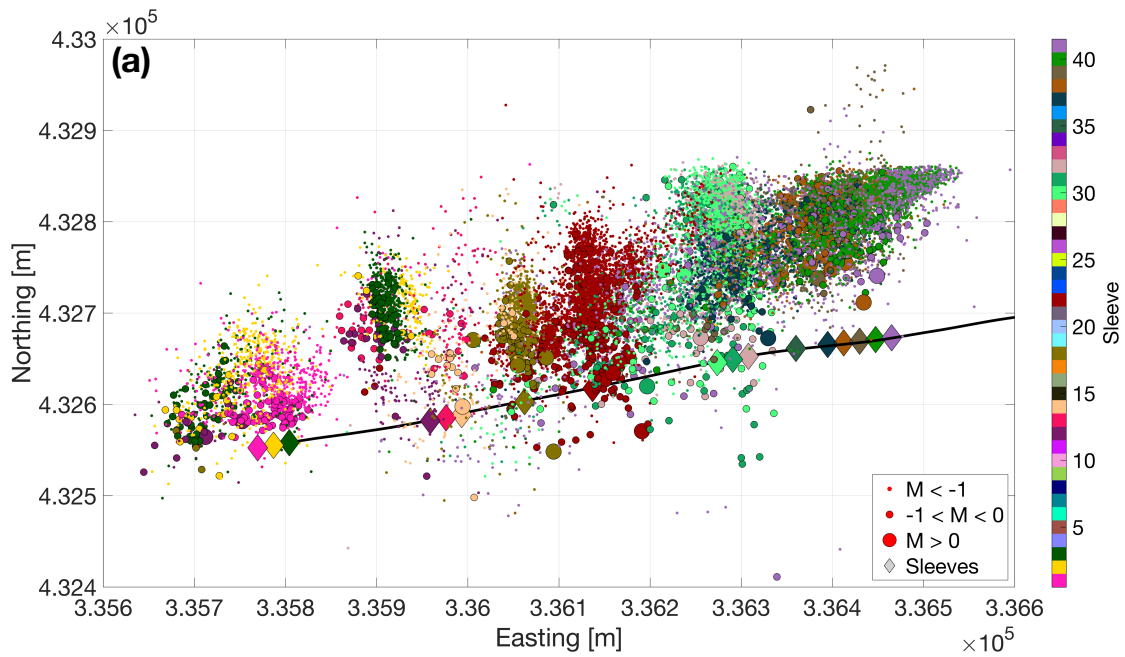
274 *Figure 4: Magnitude-frequency distribution for all events reported in real time (grey dots).*
 275 *The observed distribution follows the G-R distribution with $a = 1.9$ and $b = 1.3$ (red line). We*
 276 *use both the Wiemer and Wyss (2000) formulation (green dashed line) and a Kolmogorov-*
 277 *Smirnov test (purple dashed line) to assess the overall magnitude of completeness.*

278

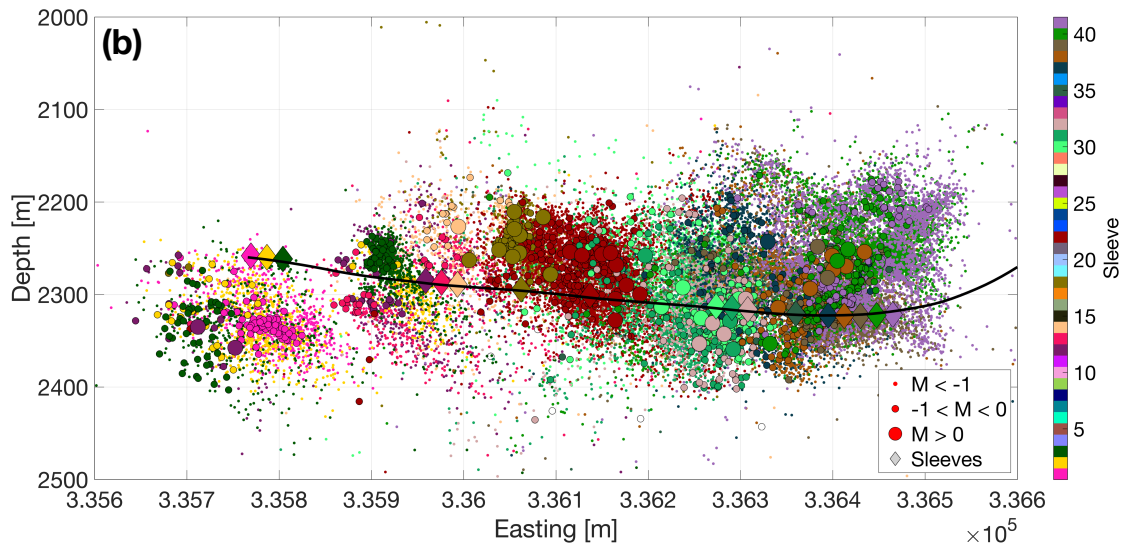
279 **3. Microseismic observations**

280 Figure 5 shows a map and cross-section for located events with a signal-to-noise ratio greater
 281 than 5. Events during each stage are mostly found in the vicinity of the corresponding
 282 injection sleeve, extending approximately 200 m to the north. The events extend
 283 approximately 150 m above and below the well, remaining within the Bowland Shale
 284 Formation. The largest observed event has a magnitude of $M = 1.5$, and in total 8 events
 285 exceeded the TLS $M = 0.5$ threshold, 3 of these occurred during injection and required
 286 pumping to be stopped, while the remaining 5 were trailing events that occurred after
 287 injection had ceased.

288



289



290

291 *Figure 5: Map view (a) and cross-section (b) of microseismic events detected during real time*
 292 *monitoring at PNR-1z. Events are coloured by the sleeve number with which they are*
 293 *associated. The PNR-1z well profile is shown by the black line.*

294

295 3.1. Relationship between microseismicity and previously-observed faults

296 Prior to the start of operations, a 3D reflection seismic survey was acquired at the site. Several
 297 pre-existing faults and “seismic discontinuities” (potential small faults that are at the limit of
 298 resolution for 3D seismic surveys) were identified (Cuadrilla Resources Ltd., 2018). We
 299 observed little or no correlation between the positions of these features the and
 300 microseismicity. The events associated with Stages 1 – 3 at the toe of the well overlap with
 301 one of the seismic discontinuities. However, the levels of microseismicity produced by these

302 stages were among the lowest. In contrast, none of the events that exceeded the $M > 0.5$ TLS
303 threshold occurred on structures identified from the 3D survey.

304 Indeed, no microseismicity coincided with any of the large faults identified in the 3D seismic
305 survey, all of which were significantly further from the well than the greatest distances
306 reached by the microseismicity. This observation allowed CRL to proceed with confidence
307 that the hydraulic stimulation was unlikely to cause re-activation of the larger faults that had
308 been identified.

309

310 **3.2. Identification of potential seismogenic structures**

311 The northwards propagation of microseismicity from each injection sleeve traces the
312 propagation of hydraulic fractures perpendicular to the minimum horizontal stress azimuth of
313 approximately 80° (Fellgett et al., 2017). However, our interest was to identify pre-existing
314 structures on which the larger events may occur. We note that the largest event, with a
315 magnitude of $M = 1.5$, could correspond to a rupture with displacement of less than 1 cm with
316 a length less than 100 m. At this scale the distinction between a “small fault” and a “large
317 fracture” is somewhat arbitrary: we will use “fault” hereafter to describe such features, while
318 keeping this fact in mind.

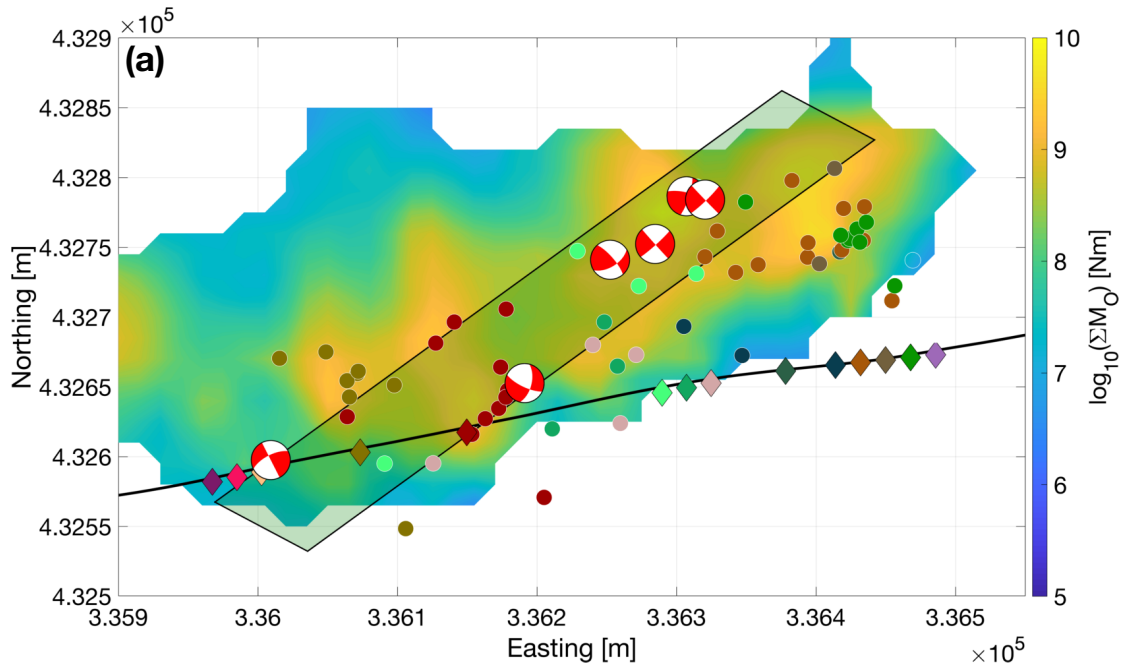
319 In Figure 5 the events do not display an obvious alignment along a pre-existing fault, an
320 observation which often provides the clearest evidence of fault reactivation (e.g., Igonin *et al.*,
321 2019; Kettlety *et al.*, 2019; Eyre *et al.*, 2019). Instead, we use a combination of observations
322 to identify and define the seismogenic structures responsible for the largest events.

323

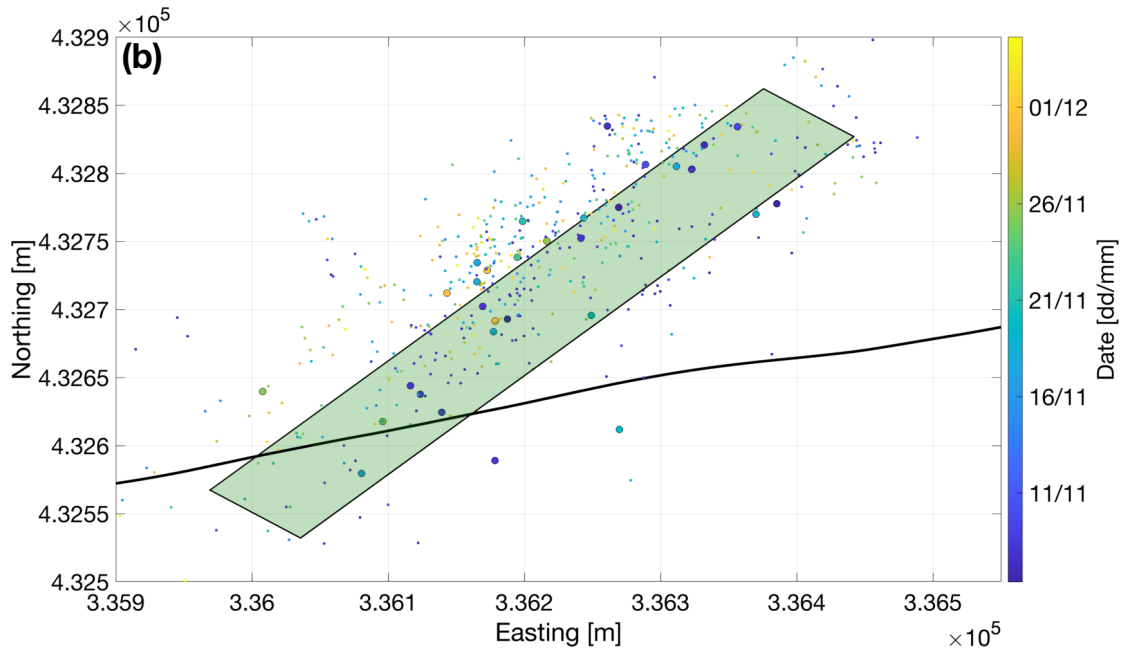
324 **3.3. Focal mechanisms**

325 The focal mechanisms for 6 of the largest events are shown in Figure 6a. The events all have
326 similar mechanisms: either left-lateral strike slip on a near-vertical fault striking NE-SW, or
327 right-lateral strike-slip on a near-vertical fault striking NW-SE. The consistent orientation of
328 these focal mechanisms provides a constraint for the orientation of any potential seismogenic
329 structure.

330



331



332

333 *Figure 6: Maps showing the observations used to identify seismogenic structures. (a) shows*
 334 *all events with $M > 0$ (dots coloured by sleeve number as per Figure 5), the cumulative*
 335 *seismic moment (contours), and the focal mechanisms of the largest events. (b) shows a map*
 336 *of the events that occurred during the injection hiatus from 3rd November to 7th December. We*
 337 *combine the largest events and the injection hiatus events to map a plane striking at 237° and*
 338 *dipping at 70° (black-outlined box).*

339

340 **3.4. Mapping large events and cumulative moment release**

341 Figure 6a also shows the positions of all events with $M > 0$, and maps the cumulative seismic
342 moment release, ΣM_0 . These observations allow us to identify a single zone in which almost
343 all of the larger events were occurring, and within which the overall cumulative seismic
344 moment release was highest. This zone intersects the PNR-1z well at roughly the position of
345 Sleeve 18, which was the first stage on which an event exceeding the $M > 0.5$ TLS threshold
346 occurred. Interaction between injection activities and this zone occurred along the well
347 towards the heel. Importantly, the orientation of this zone matches the orientation of the NE-
348 SW plane of the observed focal mechanisms.

349

350 **3.5. Microseismicity during injection hiatus**

351 These observations allowed us to identify the seismogenic feature during the initial
352 stimulation of Stages 18 – 41 in October 2018 (Figure 2). From the 3rd of November, CRL
353 paused the injection program in response to repeated $M > 0.5$ events that had occurred during
354 the previous week. The injection pause continued until 7th December. Observations of
355 microseismicity during this injection hiatus (Figure 6b) provided the final and definitive
356 identification of the seismogenic structure. The events during hiatus, almost all of which had
357 magnitudes less than $M < -1$, were all located along the same feature that we had identified
358 from the focal mechanism orientations, the positions of the largest events, and the cumulative
359 moment release map.

360 Our overall interpretation of the observed microseismicity is that a pre-existing fault plane
361 runs northeast from the well. During hydraulic stimulation, larger events occurred when the
362 hydraulic fractures from each stage intersected this fault. During the hiatus, whereas the
363 microseismic events associated with hydraulic fracturing stopped, low levels of
364 microseismicity continued to persist along this feature for a longer period of time. We fit a
365 plane to a combined population of the $M > 0$ events (Figure 6a) and the hiatus events (Figure
366 6b), by finding the plane that minimises the least-squares distance between each event and the
367 plane. We found a strike of 237° and a dip of 70° , which is consistent with the observed focal
368 mechanisms. We term this fault NEF-1 (Northeast Fault-1) hereafter. With the maximum and
369 minimum horizontal stresses oriented north-south and east-west respectively, this plane is
370 well-oriented for the observed left-lateral strike slip motion, and the observed focal
371 mechanisms are therefore consistent with the local stress conditions.

372

373 **4. Statistical Forecasting of Event Magnitudes**

374 During stimulation we applied in real time an event magnitude forecasting model to guide
 375 operational decisions with respect to induced seismicity. Hallo *et al.* (2014) introduced the
 376 concept of seismic efficiency, S_{EFF} , which describes the correlation between the cumulative
 377 moment release, ΣM_O , and the cumulative injection volume ΔV :

$$378 \quad S_{EFF} = \frac{\Sigma M_O}{\mu \Delta V}, \quad (1)$$

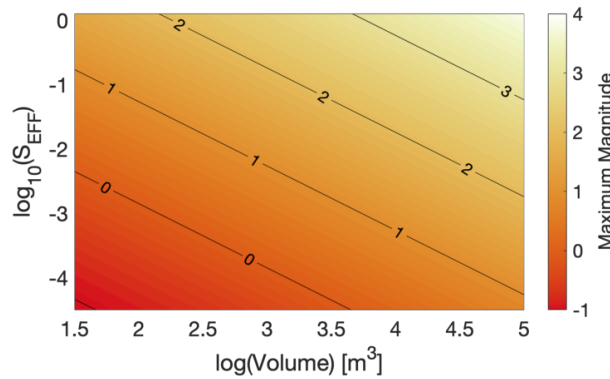
379 where μ is the shear modulus, assumed to be 20 GPa here. Based on the observed values of
 380 S_{EFF} and the b value, the size of the largest expected event, M_{MAX} can be estimated as:

$$381 \quad M_{MAX} = \frac{2}{3} \left(\log_{10} \left(\frac{S_{EFF} \mu \Delta V^{\frac{3-b}{2}}}{b 10^{9.1}} \right) \right) + \frac{2}{3} \log_{10} (10^{b\delta} - 10^{-b\delta}), \quad (2)$$

382 where δ is the probabilistic half-bin size defined around M_{MAX} (Hallo *et al.*, 2014). This
 383 formulation assumes that b and S_{EFF} do not change significantly for a given stage, or for a
 384 given volume of rock being stimulated. Verdon and Budge (2018) applied this approach in a
 385 pseudo-prospective manner to a hydraulic fracturing dataset from the Horn River Shale,
 386 Canada, showing that it would have accurately forecast event magnitudes had it been applied
 387 in real time.

388 Equation 2 posits a logarithmic dependence between injection volume and the largest event
 389 size. Given that the planned injection volumes do not vary by orders of magnitude between
 390 stages, the primary controlling factor on the largest event magnitude is therefore S_{EFF} . The
 391 relationship between S_{EFF} , ΔV , and M_{MAX} is plotted in Figure 7 (assuming $b = 1$).

392



393

394 *Figure 7: Relationship between S_{EFF} , ΔV , and M_{MAX} given by Equation 2 (assuming $b = 1$),*
 395 *showing the logarithmic dependence of M_{MAX} on ΔV .*

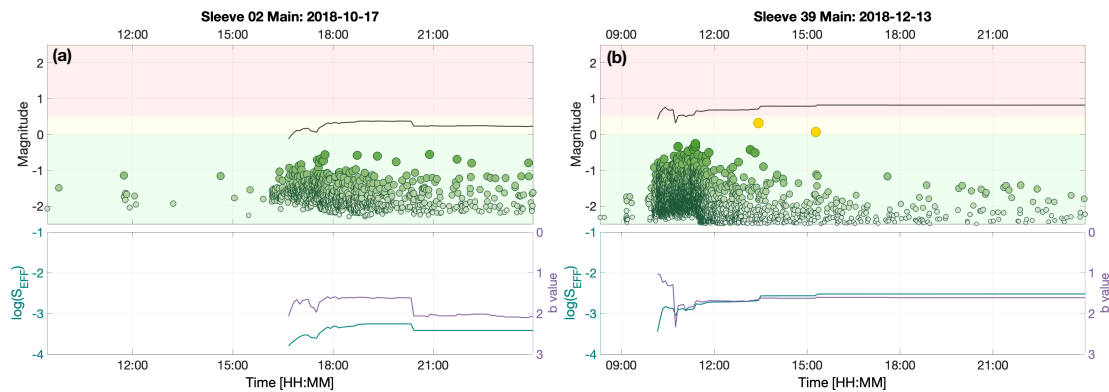
396

397 Equation (2) provides the most likely maximum event magnitude. In practice it is more useful
 398 to define a value for M_{MAX} that is unlikely to be exceeded. Using synthetic event

399 distributions, Verdon and Budge (2018) showed that adding a value of 0.5 to Equation 2 is
 400 sufficient to capture 95% of the variance between true and re-constructed model populations.
 401 In our analysis we applied this correction such that our results provided a value that, within
 402 reasonable levels of certainty, will not be exceeded.

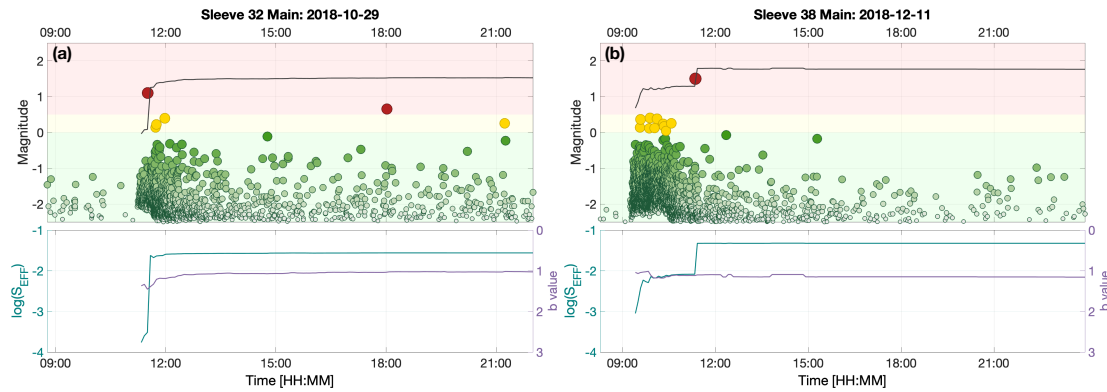
403 We tracked b and S_{EFF} in real time during every stage, providing regularly-updated forecasts
 404 of M_{MAX} . We computed the b value using the Aki (1965) maximum likelihood approach,
 405 finding the minimum completeness threshold using a Kolmogorov-Smirnov test at a 10%
 406 acceptance level to assess the quality of fit between the observed magnitude distribution and
 407 the G-R relationship (Clauset *et al.*, 2009; Williams and Le Calvez, 2013), requiring a
 408 minimum of 50 events for a reliable measurement (though with over 39,000 events in 17
 409 stages, the number of events passed this threshold very quickly for each stage).

410 Figure 8 shows a selection of results for this analysis when performed on a stage-by-stage
 411 basis, i.e., considering ΣM_O and ΔV associated with each individual stage. We find that for
 412 most of the stages this approach provided accurate bounds, with the observed events falling
 413 within the modelled value of M_{MAX} . However, this is not always the case, as can be seen for
 414 Stages 32 and 38 in Figure 9, for example.



415

416 *Figure 8: Examples of S_{EFF} , b , and M_{MAX} tracked during injection on a stage-by-stage basis.*
 417 *In the lower panels we track S_{EFF} (blue) and b (purple), and in the upper panels we plot the*
 418 *resulting values of M_{MAX} (black line) compared against observed events (circles coloured by*
 419 *magnitude relative to the TLS thresholds).*



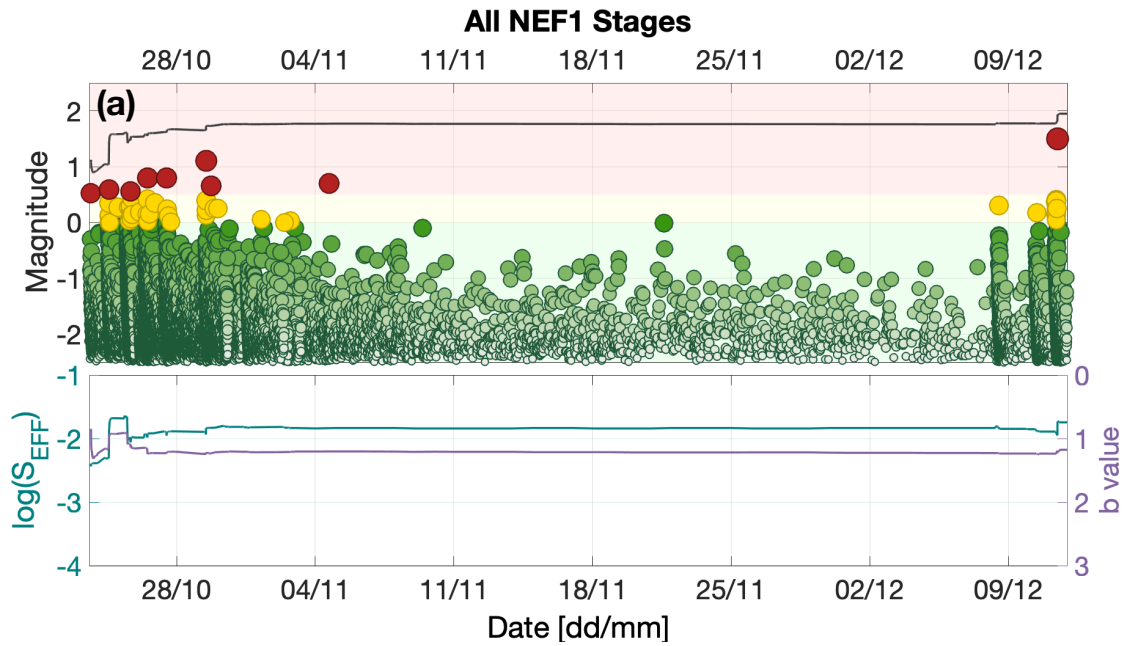
420

421 *Figure 9: Examples of S_{EFF} , b , and M_{MAX} tracked during injection on a stage-by-stage basis,*
 422 *in the same format as Figure 8. For some stages, events occur that exceed the modelled*
 423 *M_{MAX} values, when the injection volumes and observed events are treated discretely on a*
 424 *stage-by-stage basis.*

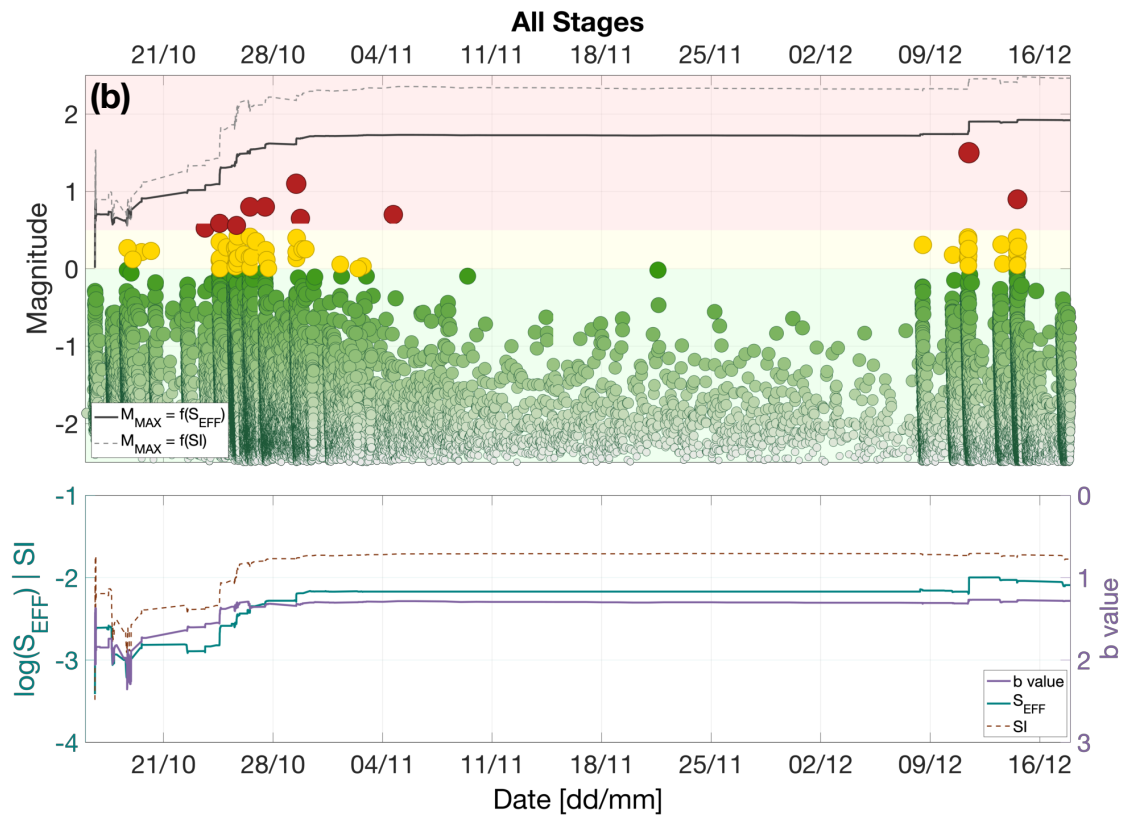
425

426 The reason for this discrepancy is obvious when considered in the light of the observations
 427 and interpretations of the microseismicity presented in Section 3: the NEF-1 fault runs
 428 obliquely to the well and was intersected by multiple stages. It is therefore not appropriate to
 429 consider each stage independently because the seismicity was caused by repeated injection
 430 into the same feature. Instead, as the NEF-1 feature was identified, we adjusted our approach
 431 to include the effects of repeated injection, treating all injection and seismicity from Stage 18
 432 onwards cumulatively (Figure 10a). The value of S_{EFF} was observed to stabilise very quickly
 433 at a value of approximately $\log_{10} S_{EFF} \approx -2$, which produces a forecast M_{MAX} of 1.7. The
 434 largest observed event at PNR-1z had a magnitude of $M = 1.5$.

435



436



437

438 *Figure 10: Forecasting M_{MAX} over cumulative stages. Here we treat stages cumulatively to*
 439 *generate M_{MAX} forecasts when (a) all of the stages that intersect the NEF-1 fault are*
 440 *considered, and (b) when all stages are considered. The observed S_{EFF} is initially at*
 441 *approximately $\log_{10} S_{EFF} \approx -3$, giving a forecast $M_{MAX} < 1$. As the injection begins to*
 442 *interact with the NEF-1 feature, the b value decreases and the overall seismic efficiency*
 443 *increases to approximately $\log_{10} S_{EFF} \approx -2$, giving a forecast $M_{MAX} < 2$. In (b) we also*

444 *show the Shapiro et al. (2010) seismogenic index (gold dashed line), and the resulting M_{MAX}*
445 *forecast from this approach (grey dashed line).*

446

447 For completeness, we also considered the cumulative impacts of the full injection volume and
448 seismicity from all the injection stages (Figure 10b). This represents the worst-case scenario if
449 all of the injected fluid was inducing events on a single seismogenic feature. Initial values for
450 S_{EFF} are low ($\log_{10} S_{EFF} \approx -3$) and b values are high ($b > 1.5$) giving $M_{MAX} < 1$. From
451 Stage 18 onwards we observed the hydraulic fracturing interact with the NEF-1 fault,
452 producing an increase in S_{EFF} to ($\log_{10} S_{EFF} \approx -2$) and a decrease in b to approximately 1.
453 This produces an increase in M_{MAX} to $M_{MAX} \approx 2$.

454

455 **5. Discussion**

456 **5.1. Operational Decision-Making**

457 The observations presented above were used by CRL to guide their operational decision-
458 making, especially during the latter injection stages in December, after the period of injection
459 hiatus in November 2018.

460 During hydraulic fracturing, placement of the proppant cannot begin until fracture breakdown
461 has occurred and fractures begin to propagate. This typically requires a minimum of
462 approximately 80 m³ of fluid. The proppant concentration is then gradually increased as the
463 injection continues, such that the majority of proppant is placed at the end of the stage. If a
464 stage is terminated mid-way through by a TLS red-light event, only a small proportion of the
465 proppant will have been placed, even if several hundred m³ of fluid has been injected. In
466 effect, the stage will therefore have been wasted and the environmental water use and seismic
467 risk unnecessarily increased.

468 At PNR-1z, the modelling described above showed that events larger than $M = 2$ were not
469 expected on the NEF-1 fault given the observed b values and seismic efficiency, and the
470 planned injection program. This forecast was reported to the OGA in November 2018, and it
471 falls within the objectives of seismicity mitigation set out by the OGA (minimising felt events
472 and avoiding damaging events). However, the NEF-1 fault could be expected to continue
473 producing $M > 0.5$ red-light events that would terminate injection, preventing the placement
474 of proppant. CRL therefore decided that further injection into the sleeves that intersect the
475 NEF-1 fault would be wasted, and in December 2018 they restarted injection in Stages 37 –
476 41 at the heel of the well. Based on the seismicity mapping described in Section 3 it was
477 hoped that these stages would pass to the east of the NEF-1 fault, allowing stages to be

478 completed without interruption. Based on the forecasting described in Section 4, CRL was
479 able to do so with confidence if these stages did intersect NEF-1, the levels of seismicity
480 would not exceed the objectives set by the OGA, and therefore injection could be conducted
481 safely.

482 In reality, some of these latter stages did intersect the NEF-1 fault, triggering two further TLS
483 events with $M > 0.5$. However, the event magnitudes remained within the levels that had been
484 forecast, as described in the section above, and within the overall regulatory objective to
485 minimise the number of felt events.

486

487 **5.2. Seismic Efficiency and Seismogenic Index**

488 The Seismogenic Index, SI (Shapiro *et al.*, 2010), is another parameter that is commonly used
489 to describe the relationship between injected volume and seismicity. Whereas the S_{EFF}
490 parameter we use here scales the injection volume to the cumulative seismic moment release,
491 the seismogenic index scales the injection volume to the number of events larger than a given
492 magnitude. Since many previous studies have provided estimates of SI , it is of interest to
493 compute this parameter for the PNR-1z dataset to facilitate a comparison. Our results are
494 shown alongside the S_{EFF} results in Figure 10, and we also plot the M_{MAX} forecasts that result
495 (at 5% probability of exceedance level) using the method described by Shapiro *et al.* (2010).
496 We note that, as found by Verdon and Budge (2018), SI follows a similar trend to $\log_{10} S_{EFF}$,
497 which is not surprising because the total moment release will depend on the number of events
498 that occur. We also find that the M_{MAX} values derived from the SI measurements are larger
499 than those derived from the S_{EFF} measurements, as also found by Verdon and Budge (2018).

500 Dinske and Shapiro (2013) catalogue SI values for a range of injection sites, finding values
501 ranging from $-9 < SI < 1$. The maximum value of SI obtained here is $SI = -1.8$, which is
502 similar to many of the geothermal projects described by Dinske and Shapiro (2013), but
503 significantly larger than those obtained for hydraulic fracturing sites at Cotton Valley (East
504 Texas) and in the Barnett Shale (Northeast Texas). However, the values obtained for PNR-1z
505 are similar to values found by Verdon and Budge (2018) for hydraulic fracturing in the Horn
506 River Basin, British Columbia, Canada, where $-4 < SI < -1$, and towards the lower end of the
507 range found by Schultz *et al.* (2018) for hydraulic fracturing sites in Alberta, Canada, where
508 $-2.5 < SI < -0.5$. The most notable past case of injection-induced seismicity in the U.K. for
509 which SI values are available is the Rosemanowes Hot Dry Rock geothermal site, for which
510 Li *et al.* (2018) found maximum values of $SI = -3.4$, significantly lower than the values found
511 for PNR-1z.

512

513 5.3. Scaling between volume and cumulative moment release

514 The underlying assumption implicit to Equation (1) is that the cumulative seismic moment
515 scales linearly with the injection volume. However, recent studies (e.g., Galis *et al.*, 2017; De
516 Barros *et al.*, 2019) have proposed alternative scaling factors, and in particular that

$$517 \quad \Sigma M_O \propto V^{\frac{3}{2}}. \quad (3)$$

518 This scaling by an exponent of 1.5 is also implicit to the Shapiro *et al.* (2010) *SI* approach,
519 since the logarithm of the seismic moment scales with $1.5 \times M_W$. Discussion continues as to
520 the most appropriate value of the scaling exponent between ΣM_O and V (e.g., Chen *et al.*,
521 2018; De Barros *et al.*, 2019).

522 In Figure 11a we track the evolution of the cumulative moment release with the cumulative
523 injection volume, and estimate a least-squares fit (in log-log space) to these data for a
524 relationship having the form

$$525 \quad \Sigma M_O = \alpha V^n. \quad (4)$$

526 Our results are shown in Figure 11a. For the overall dataset, we find a best-fit value of $n =$
527 1.6. However, it is apparent that the data may not be best described by a single value. Based
528 on our observations of which stages caused reactivation of the NEF-1 fault, combined with
529 apparent changes in slope of Figure 11a, we divide the data into 3 periods: Stages 1 – 14,
530 prior to reactivation of the NEF-1 fault; Stages 18 – 38, while reactivation of the fault was
531 taking place, and Stages 39 – 41 which appeared to miss the NEF-1 fault at the heel of the
532 well. Doing so, we find best-fit values of $n = 0.8$ for Stages 1 – 14; $n = 3.0$ for Stages 18 – 38,
533 and $n = 0.6$ for Stages 39 – 41.

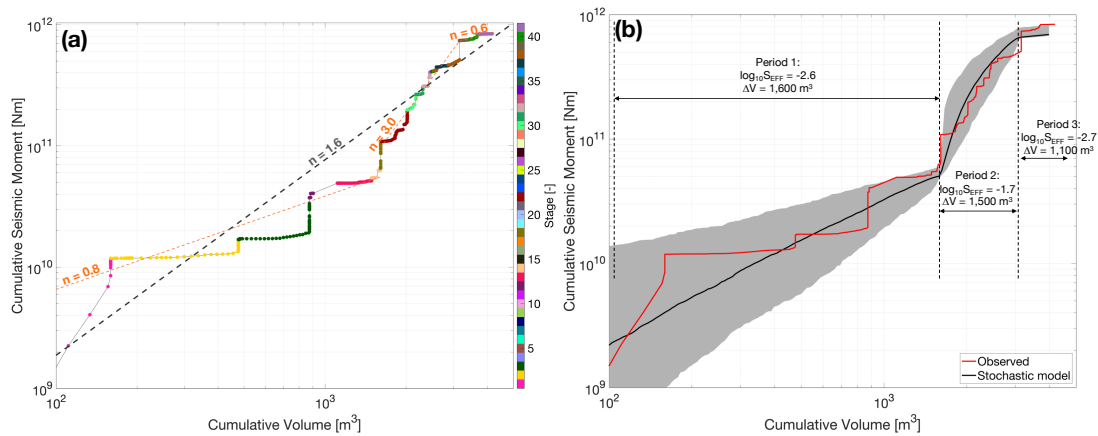
534 This variability highlights a challenge that arises when attempting to assess any scaling
535 relationship between cumulative moment and volume, should the constant of proportionality
536 (α in Equation (4)) vary during the process, which might be expected as hydraulic fracturing
537 proceeds along a horizontal well, and so encounters different volumes of rock that have
538 different geomechanical properties.

539 We further demonstrate this effect in Figure 11b. Based on our observations in Section 4, we
540 simulate a scenario whereby event populations are generated with $b = 1.2$ and $\log_{10} S_{EFF} = -$
541 2.6 (assuming a linear relationship between V and ΣM_O) for the first 1,600 m³ of injection
542 (representing Stages 1 – 14); $\log_{10} S_{EFF} = -1.7$ for the second 1,500 m³ of injection
543 (representing Stages 18 – 38); and $\log_{10} S_{EFF} = -2.7$ for the final 1,100 m³ of injection
544 (representing Stages 39 – 41). Events are generated stochastically to meet these criteria, and
545 are assumed to occur at random times within each of the specified periods. We generate 1,000
546 such populations, and in Figure 11b we plot the median value of ΣM_O as a function of V , and

547 the boundaries containing 95% of the models. The resulting models show good agreement
 548 with the observed evolution of cumulative moment release.

549 This modelling indicates the need for caution when attempting to constrain the relationship
 550 between moment release and volume: if the constant of proportionality varies during injection
 551 then a simple comparison of moment and volume may lead to under- or overestimates of the
 552 exponent n . For this dataset, a linear relationship between cumulative moment and volume,
 553 with an increase in S_{EFF} from $\log_{10} S_{EFF} = -2.6$ to $\log_{10} S_{EFF} = -1.7$ during reactivation of the
 554 NEF-1 fault, provides a good fit to the observed seismicity.

555



556

557 *Figure 11: Evolution of cumulative seismic moment with injection volume. In (a), data points*
 558 *are coloured by the corresponding stage. Power law fits to the observations are shown, with*
 559 *a best-fit exponent of $n = 1.61$ for the overall dataset (dark grey dashed line), but $n = 0.77$ for*
 560 *the early stages, $n = 2.98$ where the NEF-1 reactivates, and $n = 0.58$ during the final stages*
 561 *(orange dashed lines). In (b) the red curve shows the observed data, the black line shows the*
 562 *median stochastically-simulated model as described in Section 5.3, with the shaded region*
 563 *representing 95% of the models.*

564

565 **5.4. Assigning injection volumes to seismicity**

566 Verdon and Budge (2018) treated each hydraulic fracturing stage as an independent event,
 567 and did not treat the volumes cumulatively as injection stages proceeded. In contrast, for the
 568 PNR-1z dataset, Figures 9 and 10 show the importance of treating multiple stages in a
 569 cumulative manner, and that failure to do so would have produced a significant underestimate
 570 of the expected event magnitudes for some stages. We believe that the difference in
 571 behaviours between the two sites stems from the orientations of the faults relative to the well
 572 trajectories. In the Horn River Basin site described by Verdon and Budge (2018), the
 573 reactivated faults were orientated roughly perpendicular to the wells. As such, each

574 seismogenic feature was only affected by one or two stages (Kettley *et al.*, 2019). In contrast,
575 for PNR-1z the NEF-1 fault runs obliquely to the well, and so this feature was intersected by
576 multiple fracture stages, hence the need to treat these stages cumulatively.

577 Assigning the appropriate fluid volume when making such assessments remains a challenging
578 issue (e.g., Atkinson *et al.*, 2016). The comparison of the Horn River Basin and PNR-1z
579 examples described above shows that detailed analysis of microseismic event locations,
580 combined with a geomechanical understanding of the subsurface, is needed to guide such
581 decisions.

582

583 **6. Conclusions**

584 Recent hydraulic fracturing operations at the Preston New Road PNR-1z well were subject to
585 some of the most stringent regulations regarding induced seismicity ever applied to any kind
586 of industrial activity. The operator therefore took a proactive approach to the issue, using real
587 time microseismic monitoring to make operational decisions with respect to induced
588 seismicity. Microseismic observations allowed us to identify the presence of a pre-existing
589 structure on which elevated levels of seismicity was occurring, and to map its extent in the
590 subsurface. This structure produced multiple events that were above the TLS red light
591 threshold, forcing the operator to stop injection, resulting in wasted stages, where fluid
592 injection ceased before significant quantities of proppant could be placed. Using the
593 microseismic observations, the operator was able to move to injection locations that were less
594 likely to interact with this structure, thereby increasing the chance of conducting successful
595 stages.

596 At the same time, we used the microseismic observations to populate a statistical model to
597 estimate an upper bound for the largest expected event size during injection. This model was
598 successful in forecasting the magnitudes of the events that did occur. The forecast maximum
599 magnitudes of $M_{MAX} < 2$ was within the overall objective set by the regulator to minimise the
600 number of felt events and eliminate the possibility of damaging events. This modelling gave
601 the operator and the regulator confidence that, even if the seismogenic structure were to be
602 intersected by further fracturing stages, the level of risk posed was acceptable. This
603 confidence was borne out during operations: as further activity did occur on the identified
604 fault, but the largest event to occur had a magnitude of $M = 1.5$, within the expectations
605 provided by the statistical model.

606 Various options have been suggested to regulate induced seismicity. Fault respect distances
607 (Westwood *et al.*, 2017) require an operator to avoid known faults in the subsurface.

608 However, this case study, along with previous cases (e.g., Igonin *et al.*, 2019; Kettleby *et al.*,
609 2019) shows that reactivated faults may not be visible on 3D seismic surveys, especially if
610 they have strike slip displacement, while imaged faults may not be near to their critical stress
611 and therefore don't reactivate. Therefore the use of fault respect distances will not provide an
612 effective approach to induced seismicity regulation.

613 Whereas more advanced approaches to the mitigation of induced seismicity have been
614 proposed (e.g., Mignan *et al.*, 2017; Verdon and Budge, 2018) and demonstrated (Kwiatek *et*
615 *al.*, 2019), simple Traffic Light Schemes are the most common form of regulation applied by
616 regulators to mitigate HF-IS. The retroactive nature of these TLSs means that red light
617 thresholds may be set far lower than the actual level of seismicity that a regulator wishes to
618 prevent. Decisions are based solely on the magnitude of the largest events, which is a
619 reasonable choice if sites are monitored by regional arrays that provide limited detection
620 thresholds and poorly-constrained event locations. However, where operators acquire high-
621 quality real-time microseismic data, providing thousands of accurately-located events across
622 several orders of magnitude, then a TLS that use only the largest event magnitude, and
623 therefore discards 99.9% of the observations available, seems unnecessarily crude. In this
624 paper we have demonstrated how an operator can use microseismicity to assess the seismic
625 risk, and make proactive decisions to mitigate induced seismicity in real time. Such an
626 approach is more in line with the type of goal-setting regulation (Lindøe *et al.*, 2012) that has
627 been applied with much success to other aspects of the oil and gas industry. Induced
628 seismicity poses a risk for other forms of sub-surface industrial activity including engineered
629 geothermal systems, and the storage of CO₂ in geologic reservoirs. As induced seismicity
630 continues to attract public scrutiny, the proactive real-time use of seismic monitoring, as
631 demonstrated here, could see many other applications.

632

633 **Data and Resources**

634 The event catalogues and injection data used in this paper are scheduled to be released by the
635 Oil and Gas Authority (<https://www.ogauthority.co.uk/data-centre/>) on the 27th June 2019.

636

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645

646 **References**

- 647 Aki K., 1965. Maximum likelihood estimate of b in the formula $\log N = a - bM$ and its
648 confidence limits: Bulletin of the Earthquake Research Institute, University of Tokyo 43,
649 237-239.
- 650 Akkar S., M.A. Sandikkaya, J.J. Bommer, 2014. Empirical ground-motion models for point-
651 and extended-source crustal earthquake scenarios in Europe and the Middle East: Bulletin
652 of Earthquake Engineering 12, 359-387.
- 653 Atkinson, G.M., D.W. Eaton, H. Ghofrani, D. Walker, B. Cheadle, R. Schultz, R. Shcherbakov,
654 K. Tiampo, J. Gu, R.M. Harrington, Y. Liu, M. van der Baan, H. Kao, 2016. Hydraulic
655 fracturing and seismicity in the Western Canada Sedimentary Basin: Seismological
656 Research Letters 87, 631-647.
- 657 Baisch S., C. Koch, A. Muntendam-Bos, 2019. Traffic Light Systems: To what extent can
658 induced seismicity be controlled: Seismological Research Letters, in press.
- 659 Bao X. and D.W. Eaton, 2016. Fault activation by hydraulic fracturing in western Canada:
660 Science 354, 1406-1409.
- 661 Bosman K., A. Baig, G. Viegas, T. Urbancic, 2016. Towards an improved understanding of
662 induced seismicity associated with hydraulic fracturing: First Break 34, 61-66.
- 663 Broccardo M., A. Mignan, S. Wiemer, B. Stojadinovic, D. Giardini, 2017. Hierarchical
664 Bayesian modeling of fluid-induced seismicity: Geophysical Research Letters 44, 11,357-
665 11,367.
- 666 Butcher A., R. Lockett, J.P. Verdon, J-M. Kendall, B. Baptie, J. Wookey, 2017. Local
667 magnitude discrepancies for near-event receivers; implications for the UK traffic light
668 scheme: Bulletin of the Seismological Society of America 107, 532-541.
- 669 Cuadrilla Resources Ltd., 2018. Preston New Road 1z Hydraulic Fracture Plan. Accessed at:
670 [https://consult.environment-agency.gov.uk/onshore-oil-and-gas/information-on-cuadrillas-
671 preston-new-road-site/supporting_documents/Preston%20New%20Road%20HFP.pdf](https://consult.environment-agency.gov.uk/onshore-oil-and-gas/information-on-cuadrillas-preston-new-road-site/supporting_documents/Preston%20New%20Road%20HFP.pdf) on
672 25/3/2019.
- 673 Chambers K., J-M. Kendall, S. Brandsberg-Dahl, J. Rueda, 2010. Testing the ability of surface
674 arrays to monitor microseismic activity: Geophysical Prospecting 58, 821-830.
- 675 Chen X., J. Haffener, T.H.W. Goebel, X. Meng, Z. Peng, J.C. Chang, 2018. Temporal
676 correlation between seismic moment and injection volume for an induced earthquake
677 sequence in central Oklahoma: Journal of Geophysical Research 123, 3047-3064.
- 678 Clarke H., L. Eisner, P. Styles, P. Turner, 2014. Felt seismicity associated with shale gas
679 hydraulic fracturing: The first documented example in Europe: Geophysical Research
680 Letters 41, 8308-8314.
- 681 Clauset A., C.R. Shalizi, M.E.J. Newman, 2009. Power-law distributions in empirical data:
682 Society for Industrial and Applied Mathematics Review 51, 661-703.
- 683 De Barros L., F. Cappa, Y. Guglielmi, L. Duboeuf, J-R. Grasso, 2019. Energy of injection-
684 induced seismicity predicted from *in-situ* experiments: Nature Scientific Reports 9:4999.
- 685 Dinske C. and S. Shapiro, 2013. Seismotectonic state of reservoirs inferred from magnitude
686 distributions of fluid-induced seismicity: Journal of Seismology 17, 13-25.

687 Edwards B. and J. Douglas, 2014. Magnitude scaling of induced earthquakes: Geothermics
688 52, 132-139.

689 Eyre T.S., D.W. Eaton, M. Zecevic, D. D'Amico, D. Kolos, 2019. Microseismicity reveals
690 fault activation before M_w 4.1 hydraulic-fracturing induced earthquake: Geophysical
691 Journal International 218, 534-546.

692 Fellgett M.W., A. Kingdon, J.D.O. Williams, C.M.A. Gent, 2017. State of stress across UK
693 regions: British Geological Society, Nottingham, UK. Open File Report OR/17/048.

694 Friberg P.A., G.M. Besana-Ostman, I. Dricker, 2014. Characterisation of an earthquake
695 sequence triggered by hydraulic fracturing in Harrison County, Ohio: Seismological
696 Research Letters 85, 1295-1307.

697 Galis M., J.P. Ampuero, P.M. Mai, F. Cappa, 2017. Induced seismicity provides insight into
698 why earthquake ruptures stop: Science Advances 3, eaap7528.

699 Green C.A., P. Styles, B.J. Baptie, 2012. Preese Hall shale gas fracturing review and
700 recommendations for induced seismic mitigation. Department of Energy and Climate
701 Change, London.

702 Grigoli F., S. Cesca, A.P. Rinaldi, A. Manconi, J.A. López-Comino, J.F. Clinton, R.
703 Westaway, C. Cauzzi, T. Dahm, S. Wiemer, 2018. The November 2017 MW 5.5. Pohang
704 earthquake: A possible case of induced seismicity in South Korea: Science 360, 1003-
705 1006.

706 Gutenberg B., and C.F. Richter, 1944. Frequency of earthquakes in California: Bulletin of the
707 Seismological Society of America 34, 185-188.

708 Hallo M., I. Oprsal, L. Eisner, M.Y. Ali, 2014. Prediction of magnitude of the largest
709 potentially induced seismic event: Journal of Seismology 18, 421-431.

710 Hammack R., W. Harbert, S. Sharma, B. Stewart, R. Capo, A. Wall, A. Wells, R. Diehl, D.
711 Blaushild, J. Sams, G. Veloski, 2014. An Evaluation of Fracture Growth and Gas/Fluid
712 Migration as Horizontal Marcellus Shale Gas Wells are Hydraulically Fractured in Greene
713 County, Pennsylvania: EPA Technical Report Series, U.S. Department of Energy,
714 National Energy Technology Laboratory, Pittsburgh, PA, NETL-TRS-3-2014.

715 Hansard, 2018. House of Commons Official Report 649(208), 714-716.

716 Häring M.O., U. Schanz, F. Ladner, B.C. Dyer, 2008. Characterisation of the Basel 1 enhanced
717 geothermal system: Geothermics 37, 469-495.

718 Hofmann H., G. Zimmermann, A. Zang, K-B. Min, 2018. Cyclic soft stimulation (CSS): A new
719 fluid injection protocol and traffic light system to mitigate seismic risks of hydraulic
720 stimulation treatments: Geothermal Energy 6, 27.

721 Hofmann H., G. Zimmermann, M. Farkas, E. Huenges, A. Zang, M. Leonhardt, G. Kwiatak, P.
722 Martinez-Garzon, M. Bohnhoff, K-B. Min, P. Fokker, R. Westaway, F. Bethmann, P. Meier,
723 K.S. Yoon, J.W. Choi, T.J. Lee, K.Y. Kim, 2019. First field application of cyclic soft
724 stimulation at the Pohang enhanced geothermal system site in Korea: Geophysical Journal
725 International 217, 926-949.

726 Igonin N., J.P. Verdon, J-M. Kendall, D.W. Eaton, 2019. The importance of pre-existing
727 fracture networks for fault reactivation during hydraulic fracturing: Journal of Geophysical
728 Research, *sub judice*.

729 Kendall J-M., A. Butcher, A.L. Stork, J.P. Verdon, R. Lockett, B.J. Baptie, 2019. How big is a
730 small earthquake? Challenges in determining microseismic magnitudes: First Break 37, 51-
731 56.

732 Keranen K.M., M. Weingarten, G.A. Abers, B.A. Bekins, S. Ge, 2014. Sharp increase in central
733 Oklahoma seismicity since 2008 induced by massive wastewater injection: Science 345,
734 448-451.

735 Kettlety T., J.P. Verdon, M.J. Werner, J-M. Kendall, J. Budge, 2019. Investigating the role of
736 elastostatic stress transfer during hydraulic fracturing-induced fault reactivation:
737 Geophysical Journal International 217, 1200-1216.

738 Kwiatek G., T. Saamo, T. Ader, F. Bluemle, M. Bohnhoff, M. Chendorain, G. Dresen, P.
739 Heikkinen, I. Kukkonen, P. Leary, M. Leonhardt, P. Malin, P. Martinez-Garzon, K.
740 Passmore, P. Passmore, S. Valenzuela, C. Wollin, 2019. Controlling fluid-induced
741 seismicity during a 6.1-km-deep geothermal stimulation in Finland: *Science Advances* 5,
742 eaav7224.

743 Li X., I. Main, A. Jupe, 2018. Induced seismicity at the UK ‘hot dry rock’ test site for
744 geothermal energy production: *Geophysical Journal International* 214, 331-344.

745 Lindøe P.H., M. Baram, J. Paterson, 2012. Robust Offshore Risk Regulation – an assessment
746 of US, UK and Norwegian approaches: European Safety and Reliability Conference,
747 Helsinki.

748 Luckett R., L. Ottemoller, A. Butcher, B. Baptie, 2019. Extending local magnitude M_L to
749 short distances: *Geophysical Journal International* 216, 1145-1156.

750 Maxwell S.C., J. Shemeta, E. Campbell, D. Quirk, 2008. Microseismic deformation rate
751 monitoring: SPE Annual Technical Conference, SPE116596.

752 Maxwell S.C., M. Jones, R. Parker, S. Miong, S. Leaney, D. Dorval, D. D’Amico, J. Logel, E.
753 Anderson, K. Hammermaster, 2009. Fault activation during hydraulic fracturing: SEG
754 Annual Meeting Expanded Abstracts 28, 1552-1556.

755 Maxwell S.C., J. Rutledge, R. Jones, M. Fehler, 2010. Petroleum reservoir characterization
756 using downhole microseismic monitoring: *Geophysics* 75, A129-A137.

757 Mignan A., M. Broccardo, S. Wiemer, D. Giardini, 2017. Induced seismicity closed-form
758 traffic light system for actuarial decision-making during deep fluid injections: *Nature*
759 *Scientific Reports* 7:13607.

760 Oil and Gas Authority, 2018. Consolidated Onshore Guidance, Version 2.2. Oil and Gas
761 Authority, London. Accessed at:
762 [https://www.ogauthority.co.uk/media/4959/29112017_consolidated-onshore-guidance-](https://www.ogauthority.co.uk/media/4959/29112017_consolidated-onshore-guidance-compendium_vfinal-002.pdf)
763 [compendium_vfinal-002.pdf](https://www.ogauthority.co.uk/media/4959/29112017_consolidated-onshore-guidance-compendium_vfinal-002.pdf) on 5/04/2019.

764 Schultz R., V. Stern, M. Novakovic, G. Atkinson, Y.J. Gu, 2015. Hydraulic fracturing and the
765 Crooked Lake sequences: Insights gleaned from regional seismic networks: *Geophysical*
766 *Research Letters* 42, 2750-2758.

767 Schultz R., G. Atkinson, D.W. Eaton, Y.J. Gu, H. Kao, 2018. Hydraulic fracturing volume is
768 associated with induced earthquake productivity in the Duvernay play: *Science* 359, 304-
769 308.

770 Segall, P., 1989. Earthquakes triggered by fluid extraction: *Geology* 17, 942-946.

771 Shapiro S.A., C. Dinske, C. Langenbruch, 2010. Seismogenic index and magnitude
772 probability of earthquakes induced during reservoir fluid stimulations: *The Leading Edge*
773 29, 304-309.

774 Skoumal R.J., M.R. Brudzinski, B. S. Currie, 2015. Induced earthquakes during hydraulic
775 fracturing in Poland Township, Ohio: *Bulletin of the Seismological Society of America*
776 105, 189-197.

777 van der Elst, N.J., M.T. Page, D.A. Weiser, T.H.W. Goebel, S.M. Hosseini, 2016. Induced
778 earthquake magnitudes are as large as (statistically) expected: *Journal of Geophysical*
779 *Research* 121, 4575-4590.

780 Verdon J.P., 2014. Significance for secure CO₂ storage of earthquakes induced by fluid
781 injection: *Environmental Research Letters* 9, 064022.

782 Verdon J.P. and J. Budge, 2018. Examining the capability of statistical models to mitigate
783 induced seismicity during hydraulic fracturing of shale gas reservoirs: *Bulletin of the*
784 *Seismological Society of America* 108, 690-701

785 Wang Z. and A. Krupnick, 2013. A retrospective review of shale gas development in the
786 United States: What lead to the boom?: *Resources for the Future* DP 13-12.

787 Webster B., 2018. Cuadrilla to resume fracking seven years after tremors: The Times, 16th
788 October 2018. Accessed at: [https://www.thetimes.co.uk/article/cuadrilla-to-resume-](https://www.thetimes.co.uk/article/cuadrilla-to-resume-fracking-seven-years-after-tremors-h6lrdhxj)
789 [fracking-seven-years-after-tremors-h6lrdhxj](https://www.thetimes.co.uk/article/cuadrilla-to-resume-fracking-seven-years-after-tremors-h6lrdhxj) on 28/3/2019.

790 Wessels, S.A., A. De La Peña, M. Kratz, S. Williams-Stroud, T. Jbeili, 2011. Identifying
791 faults and fractures in unconventional reservoirs through microseismic monitoring: First
792 Break 29, 99-104.

793 Westwood R.F., S.M. Toon, P. Styles, N.J. Cassidy, 2017. Horizontal respect distance for
794 hydraulic fracturing in the vicinity of existing faults in deep geological reservoirs: a
795 review and modelling study: Geomechanics and Geophysics for Geo-Energy and Geo-
796 Resources 3, 379-391.

797 Wiemer S. and M. Wyss, 2000. Minimum magnitude of completeness in earthquake catalogs:
798 Examples from Alaska, the western United States, and Japan: Bulletin of the
799 Seismological Society of America 90, 859-869.

800 Williams M.J. and J. Le Calvez, 2013. Reconstructing frequency-magnitude statistics from
801 detection limited microseismic data: Geophysical Prospecting 61, 20-38.

802 Zang A., G. Zimmermann, H. Hofmann, O. Stephansson, K-B. Min, K.Y. Kim, 2019. How to
803 reduce fluid-injection-induced seismicity: Rock Mechanics and Rock Engineering 52, 475-
804 493.

805 Zinno R., J. Gibson, R.N. Walker Jr, R.J. Withers, 1998. Overview: Cotton Valley hydraulic
806 fracture imaging project: SEG Annual Meeting Expanded Abstracts 17, 338-341.

807

808

809

810

811