## A case study of shallow radiation fogs over CIBA: Observations and simulations (WRF and HARMONIE models)



## L. INTRODUCTION

shallow fogs) have been chosen due to the previously checked special difficulty of models simulating this type of radiation fogs.

AEMet

- clouds from Meteosat (MSG v2012.2) satellite data product (SAF NWC of EUMETSAT) to compare the simulated and observed spatial distribution of fogs.
- determine the importance of land-atmosphere processes and vegetation over shallow radiation fogs, supposed to be strongly affected by the surface.

# **3. OBSERVATIONAL DATA ANALYSIS** DAY 13 Inferred fog thickness

Figure 3. OBSERVATIONAL ANALYSIS. a) Inferred fog thickness (m) (fog considered if relative humidity > 95% at specific levels). b) Visibility information (m) from METAR reports at Villanubla airport (15km SSE from CIBA). c) Temperature (ºC). d) Relative humidity (%) (NOTE: relative humidity at 20m and 35m calculated from temperature and a fixed mixing ratio of 3.8 g/kg (approximation). e) Mixing ratio (g/kg). f) Wind speed (m/s). g) Friction velocity (m/s) from sonic anemometers.

- STRONG surface based thermal inversions (even 12<sup>o</sup> between 2m and 97m) Lower layers -> Lower temperatures -> Higher relative humidity -> FOG Higher layers-> Higher temperatures -> Lower relative humidity -> NO FOG Low values of friction velocity -> no surface cooling extended to higher levels -> no vertical extension of fogs
  - SHALLOW FOGS
- Small changes in friction velocity -> control vertical extension of fog (20-35m). Solar radiation (not shown) able to reach the ground due to shallow condition of fogs, increase friction velocity and destabilize the PBL from sunrise (no increase in wind speed). Mixing ration at lower layers mainly controlled by condensation/evaporation (formation/dissipation of fogs)
- Mixing ratio at higher layers mainly controlled by turbulent mixing.
- LLJ during nights and even during days (fogs can modified diurnal PBL)

**C. Román-Cascón<sup>(1)</sup>** (*carlosromancascon@fis.ucm.es*), G. Morales<sup>(2)</sup>, C. Yagüe<sup>(1)</sup>, M. Sastre<sup>(1)</sup>, G. Maqueda<sup>(3)</sup> and J. Calvo<sup>(2)</sup> (1) Departamento de Geofísica y Meteorología. Universidad Complutense de Madrid, Spain. (2) AEMET. (3) Departamento de Astrofísica y Ciencias de la Atmósfera. Universidad Complutense de Madrid, Spain.

Despite the well-known adverse effects of fogs over human life, its forecasting is one of the goals still not well achieved by the Numerical Weather Prediction (NWP) models. One of the reasons is because there coexist many processes acting together and affecting the fog cycle, being difficult to correctly parameterize them <sup>[1], [2]</sup>. At CIBA site (Spain) <sup>[3]</sup>, on January 2012 there was a period (3<sup>rd</sup>-15<sup>th</sup>) characterized by high pressure systems over the western of Europe, and it led to more than 10 consecutive foggy days with different features (thickness, persistence during the daytime, vertical extension, freezing temperature values...). Two of these days (characterized by strong surface-based thermal inversions with

A deep observational data analysis from CIBA (Research Centre for the Lower Atmosphere) has been carried out. This centre is located in the Montes Torozos (41º 47'N, 4º 56'W, 840 m asl), an extensive and homogeneous plateau situated over the Spanish Northern Plateau. This place is specially suitable for the development of radiation fogs during autumn and winter. METAR data from the nearly Villanubla airport (15 km SSE from CIBA) has also been used to support the existence of fogs over the zone. This airport is specially affected by fogs during winter and in fact, Villanubla means "foggy town" in Spanish. It is important to note the special hazard of shallow fogs over landing/take off maneuvers.

Weather Research and Forecasting (WRF)<sup>[4]</sup> and HARMONIE<sup>[5], [6]</sup> model outputs has been compared to CIBA observations in order to check the ability of these models simulating these shallow fogs. They have been also compared to low

As the formation of radiation fogs starts from surface, it is important to test the use of different land-surface schemes available in WRF ARW 3.4 model, including several combinations of options available with the Noah multi-physics [7], [8] land-surface scheme, a new implementation in this latest version of WRF. The effect of using MODIS land use dataset instead of the default USGS data set have also been checked. With these preliminary experiments, the authors try to



Figure 4. LWC simulations at CIBA. a) Inferred fog thickness (m). Rest) LWC simulated by HARMONIE (b) WRF-Noah (c), WRF-RUC (d), WRF-MP1 (e), WRF-MP2 (f), WRF-MP3 (g), WRF-MP4 (h) and WRF-MP2-MODIS at 10m (black), 35m (blue), 100m (red) and 300m (green). NOTE -> For a better description of simulations, see Section 2 (Models setup).

## HARMONIE does not simulate the fog during the studied period.

- WRF:
- In general, this WRF configuration simulates fog for day 13 but not for day 12
- RUC land surface scheme simulates fog for day 12 and it is better for day 13, although it tends to overestimate the vertical extension. The fog is simulated due to a higher and incorrect simulation of mixing ratio (see Sect. 4)
- WRF-MP2 (max vegetation fraction option with Noah Multi-physics land surface scheme) introduces improvements for day 12
- WRF-MP4 (less soil permeability option) shows worse results
- WRF-MP2 with MODIS land use dataset also gets worse for day 12



- DIFFICULT SIMULATION OF SHALLOW FOGS (models markedly overestimate nocturnal 2m temperature -> Underestimation of relative humidity -> No prediction of fogs)

- The use of RUC land surface scheme in WRF obtained best results in LWC BUT through a WRONG overestimation of mixing ratio - LWC amounts at 10m seems to be slightly sensible to changes in land surface processes (Noah Multi-physics) and to land use data-set.

### FUTURE WORK – 1. More experiments focusing in combinations of turbulence and land-surface parameterizations.

8. REFERENCES	[9] http
[1] Bergot T., Terradellas E., Cuxart J., Mira A., Liechti O., Mueller M. and Nielsen N. W. (2007): <i>J. Appl. Meteorol. Climatol.</i> , 46, 504.	[ <b>9</b> ] Cu>
<ul> <li>[2] Gultepe I., Tardif R., Michaelides S. C., Cernak J., Bott A., Bendix J., Müller M. D., Pagowski M., Hansen B., Ellrod G., Jacobs W., Toth G. and Cober S. G. (2007): <i>Pure Appl. Geophys.</i>, 164, 1121.</li> <li>[3] Skamarock W., Klemp J., Dudhia J., Gill D., Barker D., Duda M., Wang W. and Powers J. (2008): NCAR Technical Note.</li> <li>[4] Román-Cascón, C., Yagüe, C., Sastre, M., Maqueda, G., Salamanca, F. and Viana, S. (2012). <i>Adv. Sci. Res.</i>, 8, 11-18.</li> <li>[5] Seity, Y., P. Brousseau, S. Malardel, G. Hello, P. Bénard, F. Bouttier, C. Lac, V. Masson. (2011). <i>Mon. Wea. Rev.</i>, 139, 976–991.</li> <li>[6] Brousseau, P., Berre, L., Bouttier, F. and Desroziers, G. (2011. <i>Q.J.R. Meteorol. Soc.</i>, 137: 409–422.</li> <li>[7] Niu, GY., ZL. Yang, K. E. Mitchell, F. Chen, M. B. Ek, M. Barlage, L. Longuevergne, A. Kumar, K. Manning, D. Niyogi, E. Rosero, M. Tewari, and Y. Xia. (2011). <i>J. Geophys. Res.</i>, doi:10.1029/2010JD015139</li> </ul>	[10] Me ava [11] Mo [12] Le http://w [13] Su [14] Lir [15] Pa

## 2. SAL statistics for spatial comparisons

//www.jsg.utexas.edu/noah-mp/ art, J., Bougeault, P. and Redelsperger, J.-L. (2000). Q.J.R. Meteorol. Soc., 126: 1–30. teoFrance: The Meso-NH Atmospheric Simulation System: Scientific Documentation Part III: Physics ailable at http://mesonh.aero.obs-mip.fr/mesonh/dir\_doc/book1\_m48\_19jan2009/scidoc\_p3.pdf rcrette, J.-J. (1991). J. Geophys. Res., 96D, 9121-9132. Moigne, P. (2009), SURFEX Scientific documentation, Available at ww.hirlam.org/index.php?option=com\_docman&task=doc\_download&gid=605&Itemid=70 koriansky S., Galperinand B., Staroselsky I. (2005): Physics of Fluids, 17, 085107–1–28. Y-L., Farley R.D., Orville H.D. (1983): J. Appl. Meteorol. Climatol., 22, 1065-1092 an H.L. and Mahrt L. (1987): Boundary-Layer Meteorol 38, 185-202. irnova, T. G., Brown J.M. and Benjamin S.G. (2000): *J. Geophys. Res.*, 105 (D3), 4077-4086.



