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1. Introduction

When the matrix of a nuclear waite in porous rock dissolves in {round water, other species in the waale are ako releaaed. Some ipecies, such at the actinidea and rare-earths, areofauch low solubility that they may precipitate aa aoou aa they are dissolved, depending on the rate of matrix dissolution, their concentration in the wante **solid, their solubilities, and the ground water transport rate. We have previously studied the solubility-limited dissolution of single ipecias it controlled by the rale of advective mass transfer from the wsste surface.¹ When the waste constituents include low-solubility members of a decay chain, Ike radioactive decay daughter can form** a precipitate at the waste surface or in the rock away from the waste **surface. Here we present a simplified analysis of Ike precipitate formation and dissolution of special in a two-member decay chain. For this purpose it is assumed that flowing ground water becomes saturated with the chemical species of Ike nuclides as it flows past the waate, resulting in a constant rate of solubility-limited dissolution of the waste matrix. Assuming congruent release of other wsste con**stituents from the solid, the time-dependent amounts of precipitates **of the first and second member of the decay chain at the waste sur**face can be calculated. Numerical illustrations for the ²⁵⁴ H₂²⁵⁰ Th **chain from spent fuel are given ia this paper. Full derivations are given elsewhere.³**

2 . AaarytU

We consider one-dimensional transport Ikcougk porous rock bom sa iafaiu plane source normal to the direction of water low. Water flows unidirectionally at the same velocity throughout the **medium. Retardation ia treated by eevikbriurn sorption. Dawersionfree transport is assumed. Tke mass balance at tke interface between watte aad porous rock** *m* **skown tekemetican'y ia Figure 1.**

Tke governing equations lor transport of a two-memket decay chain are

$$
K_1 \frac{\partial N_1}{\partial t} + \mathbf{v} \frac{\partial N_1}{\partial z} + \lambda_1 K_1 N_1 = 0, \qquad z > 0, t > 0. \tag{1}
$$

$$
K_2 \frac{\partial N_2}{\partial t} + v \frac{\partial N_3}{\partial z} + \lambda_2 K_2 N_2 = \lambda_1 K_1 N_1, \qquad z > 0, i > 0 \quad (2)
$$

where subscripts 1 and 2 refer to the mother and daughter auclides **respectively, and**

 $N_i(z,t)$ is the concentration of the ith member in the water phase in **a porous medium at a distance** *z* **from the dissolution location at a time I after the start of dissolution, [kg-nuclide/m¹], s'=l,2**

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 $K_i =$ sorption retardation coefficient for the ith member, i=1,2 $\lambda_i =$ decay constant of the ith member [yr"^{*}], i=1,2

The aim of this paper is to find the mass rate of release per unit cross sectional area, ϕ_i ; the mass of precipitates per unit cross **sectional area,** *l\;* **the concentration of the nuclides at the interface** and in the field, $N_4(0,t)$, $N_i(z,t)$; and f $_{\rm Th}$, the starting time of $^{230}{\rm Th}$ **precipitation: I J,, the duration of uranium precipitate; and lj. ^k , the duration of 3M T h precipitate.**

For illustration, we assume a simplified dissolution model whereby the waste solid dissolves at a constant mass rale for a leach time T, releasing its waste constituents congruently. We assume a **waste cylinder 4 m high and 0.5 m in diameter with 1.8 Mg of heavy •>** metal, containing 1710 kg of ²³⁴U, 0.329 kg of ²³⁴U and 3.59 x 10⁻⁶ kg of ²³⁰Th initially.³ The waste is in a water-saturated repository **where the rock porosity < is 0.01 aad ground water flows at a pore velocity v of 10 m/a. We assume a leach lime T for the waste solid of 10* years and solubilities** *N'* **for uranium and thorium of 10'* kg/m³ and 10"* kg/m¹ respectively. Uranium released from the waste will initially precipitate due to the choice of solubilities.**

Fig. 1 A compartment model for the dissolution **le precipitation of a two-member decay chain**

From a mass balance in the waste we find these initial concentrations that would result in the absence of solubility limits

$$
N_n^* = 8.55 \times 10^2 \, kg/m^3 \tag{3}
$$

$$
N_4^* = 1.64 \times 10^{-8} \ kg/m^3 \tag{4}
$$

$$
N_{Th}^* = 1.80 \times 10^{-10} \text{ kg/m}^3 \tag{5}
$$

where the subscripts 4, 8 and Th stand for ²³⁴U, ²³⁴U and ²³⁰Th respectively. Because $N_4^* - N_4^* > N_U^*$, the uranium solubility, uranium **precipitates at** *I* **= 0 and continues until**

$$
t_U^* = 5.5 \times 10^9
$$
 years (6)

For the purpose of calculating the duration of the uranium precip-
itate, we include the decay of ²²⁴U, but we neglect the decay of
²³⁴U in the formation of ²³⁴U in calculating the latter's concentration. Assuming that ²³⁴U and ²³⁸U dissolve from the waste in the **same ratio as they exist in the waste, we can write equations for the concentration at the interface**

$$
N_4(0,t) = N_0^* \frac{N_4^*}{N_8^*} e^{-\lambda_4 t}, \qquad 0 < t < T \tag{7}
$$

$$
N_4(0,t) = N_0^* \frac{N_4^*}{N_4^*} e^{-\lambda_4 t + \lambda_1 (t-T)}, \qquad T < t < t_0^* \tag{8}
$$

$$
N_4(0, t) = 0, \t t > t_U^* \t (9)
$$

W,(0,l) = 0, OiJ , (9) The release function for the mother nuclide ^W U is

$$
\phi_4(t) = vN_4^e e^{-\lambda_4 t} \{h(t) - h(t-T)\}
$$
 (10)

where $h(\cdot)$ is a Heaviside step function. Eq. (10) states that, in **the absence of a solubility limit, ground water at velocity « would carry away the initial concentration of** *ntV* **exposed at the interface,** subject to decay, during the leach time. If we assume the isotopic **ratio in the precipitate is the same as the waate, we caa calculate the rate of formation of the precipitate of uranium**

$$
P_4(t) = v \frac{N_4^2}{N_4^2} (N_4^2 - N_U^2) t e^{-\lambda_4 t} \qquad 0 \le t \le T \tag{11}
$$

Eq. (11) stales the conservation of mass at the interface. The excess of the initial concentration over the solubility that is not advected away is precipitated. Here because ²²⁴U is much more abundant **than 3M U, precipitation is controlled by the concentrstion difference of ²³⁶U. Beyond the leach time, the amount of precipitate of ²³⁴U**

$$
P_4(t) = \frac{N_4^*}{N_8^*} e^{-\lambda_4 t + \lambda_6 (t-T)} P_8(t), \qquad t > T \tag{12}
$$

is directly related to the amount of precipitate of ^U *U. Assuming

$$
P_8(t) = \Psi T(N_0^b - N_U^c) e^{-\lambda_0 (t-T)} - \frac{\Psi N_U^2}{\lambda_0} (1 - e^{-\lambda_0 (t-T)}) , t > T
$$
 (13)

This completes the set of boundary conditions needed to solve the governing equation for the concentration of the mother nuclide 3M U everywhere, and a full solution can be given

$$
N_4(z,t) = N_0^2 \frac{N_0^9}{N_4^9} e^{-(\lambda_4/v_1)s} \{e^{-\lambda_4(1-t/v_1)}[h(t-\frac{z}{v_1})-h(t-T-\frac{z}{v_1})]\}
$$

+e^{-\lambda_4(t-z/v_1)+\lambda_4(t-T-s/v_1)}[h(t-T-\frac{z}{v_1})-h(t-t_U^*-\frac{z}{v_1})]\}^{7}_{\ell \geq 0} , (14)

where $v_i = v/K_U$ is the migration speed, the pore velocity divided **by the retardation coefficient of uranium.**

Now we are ready to deal with the daughter nuclide ^a °Th . At $t = 0$ ²³⁰Th is not at solubility so no precipitate of ²³⁰Th ex-
ists initially. The release rate of ²³⁰Th from the waste is also the **dissolution rale, and is**

$$
\phi_{T,h}(t) = \nu N_{T,h}^{\mathbf{e}} e^{-\lambda_{T,h}t} + \frac{\nu N_{\mathbf{e}}^{\mathbf{e}} \lambda_{\mathbf{e}}}{\lambda_{T,h} - \lambda_{\mathbf{e}}} \{e^{-\lambda_{T,h}t} - e^{-\lambda_{T,h}t}\}, 0 \le t \le T
$$
\n(15)

(15) Usiuc, il5) lo solve Uw following equation fc.T the time rrs, that precipitation of ²³⁰Th begins at the waste wurlace

$$
N_{Th}(0,t) = \frac{\phi_{Th}}{v} + \frac{\lambda_k P_4(t)}{v} = N_{Th}^* \tag{16}
$$

we get

t n **= 1088yr** The precipitation of ³³⁰Th starts well before the end of the leach **lime. To determine the duration of this aoTh precipitate we solve for a tero of**

$$
P_{TA}(t) = ve^{-\lambda \tau_{A} t} N_{A}^{\lambda} \lambda_{t}
$$

$$
\left(\left\{ 1 - \frac{N_{U}^{2}}{N_{A}^{2}} \right\} \frac{\left\{ e^{\left(\lambda \tau_{A} - \lambda_{A} \right) t} - \frac{t \tau_{A} e^{\left(\lambda \tau_{A} - \lambda_{A} \right) t \tau_{A}}}{\lambda \tau_{A} - \lambda_{A}} \right\}}{\lambda \tau_{A} - \lambda_{A}} \right)
$$

$$
- \frac{\left\{ - \frac{t \tau_{A}}{\lambda \tau_{A} - \lambda_{A}} + \frac{N_{U}^{2}}{N_{A}^{2}} \frac{e^{\left(\lambda \tau_{A} - \lambda_{A} \right) t} - e^{\left(\lambda \tau_{A} - \lambda_{A} \right) t \tau_{A}}}{\left(\lambda \tau_{A} - \lambda_{A} \right)^{2}} \right\}}{\lambda \tau_{A}} \right)
$$

$$
- \frac{\nu N_{TL}^{2}}{\lambda \tau_{A}} [1 - e^{-\lambda \tau_{A} (t - t \tau_{A})}], \qquad t \tau_{A} \le t \le T \tag{17}
$$

and obtain

$$
t_{\text{TA}}^* = 1.5 \times 10^6 \text{ yr}
$$

as the end of the ³³⁰Th precipitate. Thus the ²³⁰Th precipitate lasts **longer than the leach time and we need a modified version of eq. (17) to serve as the source term for "°Th until the end of the uranium precipitate**

$$
Pr_{\mathbf{A}}(t) = P_{\mathbf{B}}(T)e^{-\lambda_{\mathbf{F}}\mathbf{A}(t-T)} - \frac{\nu N_{\mathbf{F}}}{\lambda_{\mathbf{F}}\mathbf{A}}\left[1 - e^{-\lambda_{\mathbf{F}}\mathbf{A}(t-T)}\right] + ve^{-\lambda_{\mathbf{F}}\mathbf{A}}\left[\mathcal{N}_{\mathbf{A}}^2\mathcal{N}_{\mathbf{A}}\right]
$$

$$
\times \left(T + \frac{N_{\mathcal{U}}^2}{N_{\mathbf{A}}^2}\frac{1}{\lambda_{\mathbf{F}}\mathbf{A} - \lambda_{\mathbf{A}}}\right) \frac{e^{(\lambda_{\mathbf{F}}\mathbf{A} - \lambda_{\mathbf{B}})\mathbf{F}} - e^{(\lambda_{\mathbf{F}}\mathbf{A} - \lambda_{\mathbf{A}})^2}}{\lambda_{\mathbf{F}}\mathbf{A} - \lambda_{\mathbf{A}}}
$$

$$
-ve^{-\lambda_{\mathbf{F}}\mathbf{A}}\mathcal{N}_{\mathbf{A}}^2\lambda_{\mathbf{A}}\frac{N_{\mathcal{U}}^2}{N_{\mathbf{B}}^2}\left(\frac{te^{(\lambda_{\mathbf{F}}\mathbf{A} - \lambda_{\mathbf{A}})\mathbf{F}} - T e^{(\lambda_{\mathbf{F}}\mathbf{A} - \lambda_{\mathbf{A}})^2}}{\lambda_{\mathbf{F}}\mathbf{A} - \lambda_{\mathbf{A}}}\right), T \leq t \leq t_{\mathcal{U}}\tag{18}
$$

After t_{nk}^* , $P_{Th}(t) = 0$. Using the above we can obtain the interface concentration of ²⁵⁰Th

$$
N_{Th}(0,t)=\frac{\phi_{Th}}{v}+\frac{\lambda_4P_4(t)}{v}, \qquad 0
$$

$$
N_{Th}(0,t) = N_{Th}^*, \qquad \qquad t_2 \le t \le t_2^* \qquad (19b)
$$

$$
N_{T\Lambda}(0,t)=\frac{\lambda_4 P_4(t)}{v}, \qquad t_2^* \leq t \leq t_U^* \qquad (19c)
$$

$$
N_{T\Lambda}(0,t)=0,\qquad t\geq t_U^* \qquad (19d)
$$

Eq. (19a) slates that before the start or ""Th precipitation, the interface concentration is controlled by the rate of waste dissolution, advection, and decay from the precipitate of ²³⁴U, the mother

Fig. 2 Normalized concentrations and amounts of precipitates of the 234U-230Th chain from spent fuel, at the interface

nuclide. In the period of ²³⁰Th precipitation, the interface concentra-
tion is saturation, eq. (19b). After the ³³⁰Th precipitate has disappeared but while the uranium precipitate still remains, the interface concentration is controlled by the decay from the precipitate of the mother nuclide and advection, eq. (19c). Finally, after all precipitates have disappeared, there is no ²³⁰Th at the interface, eq. (19d).

The normalized concentrations at the interface and amounts of precipitates of ²³⁴U and ²³⁰Th are shown in Figure 2. The equations for $N_{Th}(z,t)$ are very lengthy. They are given in Reference 2 but not reproduced here. Figure 3 shows the far-field concentrations of ²³⁴U and ²³⁰Th in the critical period when there is ²³⁹Th precipitate at the interface, calculated using dispersion-free one-dimensional transport, using a previously derived analytic solution⁴ and the data in Table I. Actual numerical outputs given in Reference 2 show that 230Th concentrations in the field do not reach saturation.

	33611	23411	$m_{\rm Th}$
Decay constant, 1/a	$1.54 \times 10^{-19} 2.81 \times 10^{-9} 8.66 \times 10^{-9}$		
Sorption coefficient, cm ³ /g	æt	201	250 ^f
Retardation coefficient [®]	6.000	6.000	74,300
Migration speed, m/a	1.67 x 10 ⁻³ 1.67 x 10 ⁻² 1.35 x 10 ⁻⁴		

 ${}^{\circ}K = 1 + \rho K_D (1 - \epsilon)/\epsilon$, rock density $\rho = 3.0 \text{ g/cm}^3$ | Ref. 5.

Conclusion

We have presented a simplified analysis of the dissolution, precipitation and transport of a two-member decay chain with advection as the only transport mechanism. We show the analytic solutions for the various fluxes such as dimolution rates, precipitation rates; times

Fig. 3 Normalized concentrations of the 234U-230Th chain in rock at 10⁴, 10⁸, 10⁶ years

for the precipitate existence; and nuclide concentration fields. The numerical illustration used the $^{234}U - ^{239}Th$ chain from spent fuel.

An important limitation of this analysis is that advection was assumed to be the only mode of transport. In several other studies, we have determined that in typical U.S. repository environments, ground water velocity may be so low that diffusive-advective transport may be the determinant of dissolution rates.⁶ A more realistic model would be to include a diffusive term in the governing equations.

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