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An Analytical Comparison of Various Egyptian Soils, Clays, Shales and Some Ancient Pottery by Neutron Activation*

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The term "clay" is usually used to describe an unstructured rock composed essentially of clay minerals. The physical attributes of clays that have been of most significance to man are, of course, the plastic properties they assume upon being mixed with an appropriate amount of water which permits them to be formed into pottery. Shales are also rocks formed primarily of clay minerals, but they have been metamorphosed by geologic forces into thinly layered or laminated structures. They are usually harder than clay although they were originally argillaceous sediments that accumulated in past geologic times. Many shales when ground and mixed with water regain the plastic properties of clay and serve well for pottery formation. The Egyptian shales we have analyzed have had this property, and hence should be considered among the potential pottery making materials in Egypt.

Some clays and shales contain components other than clay minerals and are described accordingly as sandy, limey, bituminous atc. Although river alluvium and cultivated soils, which are geologically Recent sediments, usually contain a preponderance of such non-argillaceous components, they often have sufficient clay minerals to be used successfully for pottery.

The characterization of a particular clay usually involves a study of its physical and chemical properties by a variety of techniques such as X-ray diffraction, electron microscopy, thermal analyses as well as chemical analysis. Chemical composition, particularly the pattern of trace impurity concentrations, can be used along with stratigraphy and petrography, particularly the identification of fossil inclusions, to confirm attributions of different outcrops to the same geologic formations and sometimes to infer the origins of these formations.

The use of clay in pottery making antedates recorded human history, and pottery remains are a major record of past civilization. In recent years, the deductions made by archaeologists with regard to pottery classification and provenance have been supplemented by modern methods of chemical analysis such as emission spectroscopy^{1,2} and neutron activation analysis.³⁻¹⁹ The latter technique has proved to be most valuable not only because it is highly sensitive but also because it allows the simultaneous quantitative determination of a large number of elements. However, most of the investigations reported have been concerned with the examination of the articles only, and few attempts have been made previously to examine the parent clays from which the ancient pottery might have been manufactured.

The present investigation is concerned mainly with the content of trace impurities in the clays and shales in Egypt. The patterns of trace impurities and minor components may be used as a basis for classification of clays and would be useful in tracing the origin of the parent materials used in pottery manufacturing. A number of clay samples, therefore, were collected from different sites throughout Egypt, to represent all extensive natural formations from the Recent deposits down to the lower Creteceous levels, and were analyzed for sixteen elemental components by instrumental thermal neutron activation methods. Also a few specimens of ancient pottery were analyzed and compared in their minor components and trace impurities to the clays.

Clay and Shale Formations in Egypt

Geological evidence indicates that the course of the Nile through Egypt was established during the Pliocens spoch when a tectonic valley, to which the Mediterranean Sea at first had access, was formed.²⁰ Accordingly the rock formations through which the Nile valley passes are earlier than

Pliocene and the deposits occurring on the flood plain of the river are of Pliocene and later date. These later deposits are rich in clay and clay-like materials, while the surrounding earlier rocks contain considerable shales and some clays. Let us first consider the surrounding rocks through which the river flows.

At the spex of the alluvial delts the river encounters several limestone layers one of which, the Haadi formation, contains significant shale intercalations. This formation extends roughly from Cairo to slightly south of the Fayum. Further south in the region of Upper Egypt which extends roughly from Qens to Idfu the river valley cuts down through the upper relatively pure limestone layers into a series of more argillaceous layers many of which are rich in shale. These shele-rich layers can extend in all to a depth of from one to two hundred meters. The layers in this region which are particularly rich in shale are in order of increasing depth the Esna shales, the Dakhla shales, the Phosphete formation and the Quseir Variegated shales. They range in date of deposition from the upper Faleocene through the upper Cretaceous. From roughly Idfu to Aswan the river valley penetrates yet deeper into the Nubia Sandstone layer which underlies these shale beds and which dates to the lower Cretaceous. The Nubia sandstone layer which is surface exposed in the Aswan region contains clay deposits which are used today for pottery manufacture.

Three distinctly different deposits lie on the flood plain of the Nile. The lowest lying of these layers accumulated during the Pliocene. It is thought to have been deposited under marine or estuarine conditions. Marine penetration into the valley had apparently ceased to be a major factor by the time of the Pleistocene epoch. During the Pleistocene North Africa was not arid as it is today but experienced periods of

abundant rainfall.²¹ Accordingly although today the tributary wadds of Egypt contribute practically no mediments to the Nile flood plain, during the Pleistocene epoch the wadie were relatively active, contributing mignificantly to the Nile mediments. The pluvial periods of the Pleistocene produced a distinctly measure second layer of deposition upon the Nile flood plain, which, as the analyses reveal, is distinctly different in chemical composition from the Pliocens layer which it covers. However it is basically similar in composition to the shale layers through which many of the contributing wadis cut. It would seem logical that the eromion of these shales contributed significantly to the formation of the Pleistocene clays.

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The upper layer of Nile alluvium has been deposited during the Recent epoch and represents predominantly foreign sedimentary material carried by the Nile from Ethiopia and East Central Africs. The analyses show it to be distinctly different in composition from all other materials studied.

The occurrence of many of these formations and deposits throughout much of Upper Egypt is shown in Fig. 1. In this figure one sees that the recent alluvium deposit lies closest to the river with a strip of Pleistocene deposit usually occurring at its outer edges. Occasional areas of Pliocene deposition surface at the extreme outer edges of the valley beyond the Pleistocene strips. The surrounding outcrops of shale are particularly evident in the region shown in Fig. 1, as is the very low lying Nubia sandstone which particularly becomes a surface rock in the Aswan region.

In the great depressions of the Western Desert in which the major oases lie, many of the same formations, i.e., the lower lying shales and Nubia sandstone, surface. Also there occur at the cases Pliceene,

Pleistocene and Recent deposits analogous to those in the Nile valley. A number of our specimens have come from these cases sites.

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There are in addition some relatively pure kaolin clays which occur in Carbouiferous Ataqa sands and Mubia sandstone on either side of the central-upper region of the Gulf of Suez and also underlie the Nubia sandstone at Aswan.

The Sources and Compositions of the Geological Specimens

The specimens of clays and shales we have analyzed were provided from the collection of the Geological Museum at Cairo through the kind permission of Dr. R. Said and the cooperation of Nr. A. Talaat. Dr. Said also generously advised and assisted in the provision of some of the samples for analysis. For the most part four or five samples of each formation considered in the previous discussion were ' ulyzed. The samples of each formation were selected to represent, as a rule, geographically disperse sites in order that the compositional uniformity of the formatione could be estimated. The source locations for the specimens analyzed are all indicated in Fig. 2.

We recognize clearly that such limited sampling as we presently have cannot be regarded as being comprehensive of the extensive formations being studied. However, particularly because the different samples of the same formations have tended to be similar in composition, this sampling does serve to indicate in general terms the basic minor and trace component compositional characteristics of these various formations. We very much hope it will be possible to add to and refine these data.

For simplicity of presentation we shall consider together the specific geographic sources of the specimens and the analytical results obtained on them, leaving for later a description of the method of analysis. Notations

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of the source locations and numerical analytical data for the individual specimens are listed together in Tables I, III, IV, and V.

Recent, Pleistocene and Pliocene Clays and Soils The analyses of the clay and soil specimens taken from Recent, Pleistocene and Pliocene deposits are grouped together in Table I. Specimens of Nile alluvium came from the Boulac section of Cairo, an island in the river at Assicut and geveral sites in and about Aswan. These alluvium specimens from three quits widely separated regions along the Nile are very closely similar in composition. Our results are also in close agreement with similar analyses of Nile alluvium from upper Egypt published by Perlmon and Asaro.^{16,13} Nile alluvium is outstanding in its very high manganese, iron and cobalt contents, as compared with all other clays analyzed. The average manganese concentration is about tenfold that in any of the other clays and the cobalt concentration is about twice as great. Conversely the thorium concentration in the alluvium is significantly low, being often about half that of the other clays and of most shales. These elements are particularly diagnostic for this material. The sodium content is also significantly high in alluvium but it is similarly high in Pleistocene clays and the shales which were found near the river.

Two other specimens of recent soils were analyzed. One came from the village of Dakhla at the Dakhla Oasis in the central Western Desert, and the second was from the Wadi El Natrun in the Western Desert south of Alexandria. The data on these specimens in Table I show that the Dakhla Oasis soil has all of the unusual compositional characteristics of the Nile alluvium. It is sufficiently similar to the alluvium in the concentrations of all components determined to suggest a common origin for the two deposits, and hence a probable early water connection between the river and

the casis. In contrast to this the soil from the Wadi Natrun does not have either the high manganese and iron concentrations or the relatively low thorium concentration that characterize Nile alluvium. Hence a different origin for this deposit is indicated.

A visual comparison is made in Figs. 3 and 4 between the standard deviation ranges of concentration of eleven of the components determined as they were found to exist in the different types of clay like material we have analyzed. One observes, for example, at the right of Fig. 3 how far the manganese content of the matching recent soils lies outside of the statistically probable ranges of its occurrence in other materials studied. Some of the components not plotted show similar separations. The only reason for not including them in the plots was to keep the visual display within modest and easily comprehensible limits.

The Pleistocene clay samples came from two regions along or near the Nile. The first region lies just south of Cairo where we have a sample from the village of Magdi and another from the Fayum. The second region is central Upper Egypt where the clay specimens came from Qena and Esna. It is worth noting that the Pleistocene clays of Qena and Esna are used extensively today for the manufacture of pottery.

Comparison of the analyzed results on these four specimens shows that although all four are basically similar in composition, the pair of specimens from south of Cairo show some consistent differences from the pair of specimens from Upper Egypt. Specifically they have more iron, cerium, europium and thorium and less barium and chromium than the clays from Qens and Esns. We have found these same differences to occur between the shales that outcrop in the region south of Cairo, i.e., the Mgadi formation shales, and shales that outcrop in the vicinity of Qens and

Esna, the Esna shales. Table II displays side by side the average concentrations and group standard deviation ranges^{*} of the compositions we have found in both

For simplicity of presentation the arithmetic means and standard deviations have been listed in Tables II and VII. Experience has convinced us that trace impurity concentrations tend to be logarithmetically distributed and hence geometric means and standard deviations in terms of logarithms of concentrations would be a more correct presentation of the data. However, because the differences between the two methods of data handling are small when the members of a group are reasonably closely similar to each other it is reasonable and practical to use the considerably more familiar linear data presentation in these instances.

the clay and shale samples from Maadi and the Fayum and in the clay and shale samples from Upper Egypt. In Table II the components whose concentration ranges are different in the two regions are underlined. It can be seen that upon the basis of these data the Pleistocene clays and Maadi formation shales from Maadi and the Fayum are indistinguishable from each other. The Pleistocene clay and Esna shale of the Qena-Esna region are not quite as closely matched in composition but are still very similar to each other. Considering the small number of samples analyzed it would be stretching statistics to distinguish between them. However, the Maadi formation shales show the same compositional differences from the Esna shales that Maedi Pleistocene clay does from Esna Pleistocene clay. Indeed the local clays of Pleistocene deposition do seem to have been derived to a predominant extent from the shales local to them.

This conformity to local conditions suggests that in other localities throughout Egypt the composition of Pleistocene clays might be characteristically different. For example, Pleistocene deposits occur along the edges of the

delta. At Alexandria, situated at the western edge of the delta, there should be available Pleistocene clays as well as alluvium for pottery fabrication. It will be necessary to characterize these and other local Pleistocene clays not yet analyzed before one can consider this study to be at all complete.

Clays of the Pliocene epoch are widely distributed over the Nile Valley and the cases. They are commonly called "plastic clays" and are used for pottery making. The four specimens analyzed came from widely separated locations. The first was from the site of the Fom 21 Khalig which was a canal within Cairo with an entrance to the Nile just east of the island of Roda. The second Nile specimen comes from the former village of Korosco, which was located south at Arwan well into Lower Nubia. The remaining two specimens came from the Central Western Desort oases, Kharga and Baharia. These four specimens are all similar to each other in composition. It would seem probable that they were all formed under similar circumstances, presumably marine or estuarine conditions. Considering such conditions of formation it would seen more likely that the Pliocene deposits would be more generally uniform than the Pleistocene and hence that it is unlikely Pliocene clay from particular localities throughout Egypt can be differentiated upon the basis of their minor and trace impurity patterns.

The Pliocene clays are characterized by relatively high concentrations of cerium, hafnium and thorium and low concentrations of sodium, iron and cobalt, characteristics which they share with the kaolin clays. In general these two types of clay are very similar in the concentrations measured. However, the kaolins have significantly less potassium, iron and manganese than the Pliocene clays, so that the two are easily differentiated upon the basis of these elements.

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Cretaceous Clays

Both the kaolin clays and Nubia sandstone clays are attributed to the Cretaceous period. The data on these clays are listed in Table III. We have analyzed three kaolin clay specimens, one from Kalabsha, south of Aswan in Lower Nubia and two from Abu Darag, near the head of the Gulf of Suez on its west coast. The specimens are reasonably similar to each other in composition, and at this stage we are considering them only as a group. Analysis of additional samples might provide some basis for differentiating the Nubian from the Suez Gulf kaolins.

As has already been noted, the kaolins are compositionally similar to the Pliocene clays but contain significantly less potassium, iron and manganese. They are characterized by their unusually low concentrations of these elements and of cobalt and by relatively high concentrations of cerium, hafnium and thorium.

All of our five Nubia sandstaone clay specimens come from the vicinity of Aswan. They include a group of three samples taken from an exploratory shaft near Aswan. They were selected for analyses in order to obtain an indication as to whether any significant changes would occur in concentrations as a function of depth from the surface. Recognizing that clays can undergo chemical exchange and also that a deposition or leaching out might occur at the surface of a clay deposit one feels that superficial samples, such as most of our specimens were, might differ from subsurface samples. There might also be natural nonuniformity throughout the deposit of course. The three specimens, T 21, T22, and T 23, taken respectively at the top layer of the clay deposit, near the shaft opening, and seventy-five meters within the shaft show a consistent change with distance in sodium, manganese, cobalt, and perhaps iron. Of these the

relative change in sodium concentration was by far the most serious one. This confirms a general feeling one has that sodium would be among the least dependable of the components measured. Chemically, sodium and other alkali metal ions would be among the most labile toward chemical exchange and most likely to be present in soluble, and hence transferable, salts. It is interesting to note that the elements which show a diminution from the surface inward are ones which are present in significantly higher concentrations in the Nile alluvium than in the clay.

Despite the differences just noted the Nubia sandstone clays form a reasonably consistent group which is compositionally different from the others studied. The difference resides in the overall pattern of constituents, there being no single component present at a concentration level significantly different that is from all other clays and shales analyzed.

Shales from Near the Nile

The Maadi formation, the Esna, the Dakhla, and the Phosphate formation shales were all specimened at sites near to the Nile. As has been mentioned, the Maadi formation shale specimens came from Maadi and the Fayum. The three other shale containing layers were sampled in the El Sharawna District between latitudes 25° and 26° and longitudes 32° and 33° which is the central region of Fig. 1. Additional specimens from the Phosphate formation were taken from the Gebel Abu Had, the site of which is shown in the upper right corner of Fig. 1. Analytical data for all of these specimens are listed in Table IV.

The small but significant differences between the Maadi formation and the Esna shales were discussed when similar differences in the Pleistocene clays associated with them were considered. The Dakhla shale is little

different from the Esna shale in composition, having perhaps somewhat less iron, manganese, and cobalt. The Phosphate formation is not a uniform layer and the shales from it show considerable compositional variation. All of these shales are roughly alike in composition, however, and can be distinguished compositionally from all the other clay-like materials analyzed except the Pleistocene clays.

Shales from the Kharga Oasis

The shale bearing layers that outcrop in the Sharawna District also outcrop at the Kharge Oasis. From the Gebel Abu Tartour there we have received specimens of Dakhla shale, Phosphate formation shale and of the yet lower lying Qusier variegated shales. Data for these specimens are listed in Table V. The Dakhla and Phosphate formation shales from the casis are similar to those from the Nile with two very pronounced exceptions, both of which occur in the alkal metal elements. On the average there is only about one-twentieth as much sodium in the shales from the casis as was in the corresponding shales from the Sharawna District and less than half as much cesium. Similar low concentrations of sodium and cesium are found in the variegated shale specimens from the Kharga Oasis. Whether these differences in alkali concentrations between shales of the two regions represent surface alterations of the type previously discussed or are basic to the shales and represent consistent differences in their conditions of deposition can be answered only through further investigation. It would be most helpful to have subsurface specimens of these shales for comparative analysis. Also it would be interesting and worthwhile to know if the Pleistocene clays at the casis also have the low sodium and cesium of the shales which outcrop there. However this may be, surface shale material taken from the casis can easily be distinguished from the other shales studied upon the basis of these elements.

In other respects the variegated shale specimens from the Kharga Oasis differed significantly from the other shales studied only in having a higher potassium content on the average than all other shales and as high an iron content as the Maadi formation shales

Archaeological Evaluation of the Clay and Shale Analyses

The analytical data on Egyptian pottery forming materials is just extensive enough at this stage to permit a reasonable evaluation of the potential of investigating ancient Egyptian pottery by means of comparative analyses. However, in order for extensive reliable inferences to be drawn from comparison of specific pottery compositions to that of the clays, the cataloging of clay compositions must be considerably more comprehensive and extensive than it now is. Also there is now need for a systematic analytical survey of ancient Egyption pottery itself to determine to what extent the various clays were used by themselves to produce pottery or in various combinations with each other and with the addition of tempering materials.

At present it looks as if it might be impossible to obtain an indication of the geographic source within Egypt of pottery made of pure Nile alluvium or Pliocene clay or mixtures of the two unless the analysis of a sufficient amount of pottery of known date and geographic provenance establishes that traditions of the use of these pure clays or particular mixtures of them were confined to specific localities at specific periods. The picture is much more encouraging with respect to other clays and the shales as there does appear to be some local compositional differentiation within the Pleistocene clays and shales. Also some of the clays which can be identified as to type through analysis, such as the Nubia sandstone and the kaolin clays, outcrop in only limited regions throughout Egypt.

It would seem likely that their use in pottery would have been similarly localized.

At this stage we have a better feeling for the consistency and diversity of clay compositions than of ancient Egyptian pottery compositions. A logical step at this time would be the undertaking of a reasonably comprehensive program of analysis of pottery of at least one historic period from known locations throughout Egypt. In this way hopefully an estimate of the consistency and degree of localization of the traditions of use of clays and tempers in pottery formation can be obtained. The insight that might be gained into the technology of the ancient Egyptian potters would itself justify such a study.

In summary, therefore, the fact that it now appears that one can distinguish between different types of clays that occur widely throughout Egypt through neutron activation analysis and in some instances one can identify clays that are local to specific regions establishes that there is potentially a reasonable likelihood that similar analyses of ancient pottery should in some instances provide indications of the geographic origins of that pottery. Because of the wide geographic distribution of some clays of nearly uniform composition it is to be expected that in many instances it may not be possible to infer such origins from the analytical data. As has been argued, the axtent to which this is true will depend greatly upon the extent to which unique traditions of pottery formulation were observed in specific locations. Only analysis of the pottery itself can answer these questions. In any event it should, in many instances, be possible to interpret the compositions of the analyzed pottery in terms of the clays from which it might have been formed.

Analysis of Pottery

A small group of pottery specimens were analyzed as a very preliminary test of the potential of correlation between clay and pottery compositions. Only nine specimens in all were studied but they were all excavated material of known origin. They were selected to range in date from early Old Kingdom into the Islamic Period and they were supplied to us through the kind cooperation of Dr. Zaky Iskander.

Three of the specimens were from Sakkara and differed greatly in their date of burial. Two were of the Old Kingdom, one, Pl, attributed to the Archaic period (32nd-28th centuries B.C.), the other, P9, attributed to the Pyramid Period (2723-2242 B.C.). The third Sakkara specimen, F8, was attributed to the relatively late 26th Dynasty (663-525 B.C.). Four specimens were of the Ptolemaic Period and were all from the Hadra excavations which were carried out at Alexandria in 1971. All four specimens P4, P5, P6, and P7 have been attributed to 140 B.C. Two specimens of yet later date, P3 and P2, were from Kom el-Dekka, also at Alexandria, and were respectively of the Coptic and Islamic periods. The data for these pottery specimens are given in Table VI.

The specimens can be divided rather well into but three compositional groups, a fact that lends credence to the concept that consistent traditions in pottery formulation did exist. The first compositional group includes both the oldest and youngest specimens which are apparated by a time span of as much as four thousand years. This is not too surprising, however, as the composition of this particular group is essentially that of the pure Nile alluvium. A comparison between the average concentrations in this group of pottery and that in Nile alluvium is presented in Table VII. One can see that they differ significantly only in the sodium concentration

which is somewhat higher in the ancient pottery. However, sodium chloride is known to accumulate as efflorescent deposits on the walls of tombs in Egypt and it is quite possible that some extra sodium could have sccumulated in the pottery during burial. It is also possible, of course, that alluvium is slightly less rich in sodium today than it was millennia ago. In all

other elements the compositions of these specimens are as close to that of Nile alluvium as one would expect them to be if indeed they had been made totally of this material. The specimens include all of the three from Sakkara and the Islamic specimen from Kom el Dekka at Alexandria. Perlman and Asaro¹⁸ have also reported on a pottery group from a cemetery at Ballas in Upper Egypt which closely matches Nile alluvium.

The second group of matching specimens include a Ptolemaic specimen from the Hadra excavation and the Coptic specimen from Kom el Dekka, both at Alexandria. These specimens have many of the characteristics of Nile slluvium, e.g., relatively high manganese and cobalt and low thorium, but to a much less degree than the alluvium itself. It seems probable that this pottery contained Nile alluvium mixed significantly with one or more other components. We have not as yet been able to match this composition with a composite of the clays we have analyzed. It is possible the material we presume to have diluted the alluvium might be a tempering material or a clay, such as a Pleistocene clay unique to the delta region which we have not yet analyzed.

The third group includes three specimens, all of Fcolemaic date from the Hadra excavations at Alexandria. Two of these three specimens are closely matched in composition and have the characteristics of Nile alluvium diluted with other components. The third is more dilute then the first two

in components that tend to be high in alluvium and could well be similarly constituted but with a lower alluvium content. All three are characterized by having definitely higher chromium than we have encountered in other materials studied except for a few of the very deeply lying shales.

These few analyses of pottery like the clay analyses encourage one to believe that more extensive studies of Egyptian pottery may often provide a basis for worthwhile inference but sometimes may only confirm an Egyptian origin for the objects studied. One considers future programs of neutron activation analysis of ancient Egyptian pottery with reserved optimism.

APPENDIX

Analytical Procedure

(1) Sampling: To insure representative sampling, at least 0.5 gram samples were ground and thoroughly mixed. In turn 40 mg samples of this more uniform material were packaged for activation. In the case of pottery, the surface of the sherd was cleaned with a tungsten carbide burr and a drill of tungsten carbide whose base could be placed in the chuck of an ordinary electric hand drill. The first portion of the powder drilled was discarded and then more powder removed to provide the amount required.

(11) Irradiation: Activations were carried out at the Brookhaven High Flux Beam Reactor at fluxes of 1-4 x 10^{14} neutrons cm⁻² sec⁻¹. A short activation of a few minutes preceded measurement of the short lived activities of manganese, sodium, potassium, and sometimes lanthanum, while an activation lasting for some hours followed by a decay period of a week to ten days preceded measurements of the longer lived activities. With every batch of samples each of the six U.S. Geological Survey Rock Standards were introduced to be irradiated under the same conditions. In this way each element is individually calibrated in each run. Agreement between

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the six standards confirms that conditions of activation and counting are uniform.

After irradiation and cooling, the samples and the standards were transferred quantitatively, using ethanol, into glass tubes provided with metal caps to allow them to be picked up by the magnet tipped arm of an automatic sample changer. An alternative more convenient procedure which does not involve the transfer of the activated samples is to use small sealed quartz tubes for packaging the samples instead of the usual aluminum cans. The main impurities in the quartz were Sb, Sc, Ca, Th, Cr, Fe, and Co, but their concentrations were negligible as compared with those encountered in the samples. Therefore the samples could be counted while still remaining in the pure quartz tubes.

(111) Counting and Spectrum analysis: The gamma ray spectra were taken with a 35 cm³ Ge(L1) detector coupled to a 3200 channel analyzer with magnetic tape output. An automatic sample changer would introduce the samples into a well in front of the detector where they were positioned in an exact geometry relative to the detector for the prescribed counting intervals. Counts from a precision pulser were fed through the entire electronics of the counting system during the period of counting, and recorded as an additional peak in the spectrum. The ratio of the number of counts collected in this pulser peak to the number of pulses emitted by the pulser during the counting period provided an exact measure of the percent live time of the entire electronic system.

(iv) Data processing: The data on the magnetic tape were processed by a Control Data 6600 computer in successive steps that 1) analyzed the spectra, 2) calculated calibration constants from the data on the six standards run with the samples, 3) converted peak data to concentrations

and 4) produced comparative plote of specimen data or carried out statistical analysis on the data for groups of specimens as desired.

These analytical procedures have been described in detail by Abascal, Harbottle and Sayre.²²

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TABLE I

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CLAYS AND SOILS DEPOSITED AFTER THE PORMATION OF THE HILE

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pech of ta	pesition				RECENT					PLEIS	TOCENE			PLIC	DIE	
Kontee Lorations		tioulac	telands	South of Asvan	Azvan	Agha= Khan Tomb Asvan	Dakkla Dakkla Gasis	Vadi Natrus	Handi South- anst of Cairo	Qeni Upper Egypt	Kona Upper Egypt	81 Fayun	Tom El Khalig Caire	Korasco Nubie	Kharga Eharga Gasis	Bahari Gasis
			ف المالية مناور مدر	Hatchi	n <u>g</u>			metching								
onclaen %	••	12	T7	T24	725	T26	715	T1	TS	T#	T9	T14	T3	711	T16	<i>i</i> 17
÷									Concentrations of	Oxides in						
od i un	nazet –	1.22	1.26	1.44	1.51	1.65	1.55	0.59	1.52	1.16	1.94	1.80	0.190	0.084	0.065	0.09
diabsiut	^K 2 ¹⁷	1.21	1.33	1.55	1.24	1.29	1.25	0.52	1.53	1.06	1.07	1.30	1.35	1.13	1.02	0.65
ron	Fe 2 ^G 3	11.3	10.8	11.4	12.2	11.4	10.1	6.9	8.16	6.39	6.64	7.68	1.43	1.90	1.80	2.24
									Concentrations of Oxid	in in Part	ts Par Hi	11100				
abidios	ah ju	76	44	43	46	41	109	50	78	55	74	45	35	31	21	30
ars i um	65.0	0.7	1.0	1.4	1.8	1.5	1.4	2.7	4.7	3.8	5.7	3.6	1.3	1.9	1.3	3.0
ar i dh	444	510	460	540	520	520	620	640	380	450	720	240	560			330
•canuium	sejū,	41	4 Û	40	40	38	38	46	55	29	33	36	30	36	36	36
an than wa	ر 0 ر هنا	58		42	37	36	35		62	107	68	55	88	82	95	65
Coffin	Cev,	150	138	90	84	78	145	155	144	47	73	183	480	440	500	320
Europium	Eu ₂ 01	2.7	2.9	3.1	2.8	2.9	2.5	3.0	3.1	2.0	2.9	2.8	2.7	2.8	3.8	2.5
Haffilum	HEO,	6.8	6.9	8.5	6.8	7.6	7.6	7.3	5.9	3.2	3.1	7.4	22.6	19.8	22.6	15.3
Thorium	th0,	7.7	6.5	7.0	6.2	6.0	6.4	12.3	17.5	13.7	14.6	13.7	21.0	25.5	26.3	16.4
lactalim	•	3.5	3.7	2.2	2.1	2.5	2.9	4.8	3.7	2.4	2.8	3.9	3.0	4.7	3.6	5.
dhreniun	• •	203	195	250	218	235	223	195	194	245	281	208	157	176	175	171
Masyunese	~ /	1740	1840	1660	1520	1540	1520	104	169	295	165	112	102	44	90	139
		./40	1940	1444	1714	2.770	1.)LV		474	474	147	-1-			15	

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		Specimens fr	on Nile Region	Specimens from the Qena-Esna Region of Upper Egypt			
		South o	f Cairo				
		Pleistocene Clay	Meadi Formation Shale	Pleistocene Clay	Esna Shale		
			Average Concentra	tions in Percent*			
dium	Na ₂ 0	1.66 ± 0.20	1.70 ± 0.05	1.55 ± 0.55	2.43 ± 0.25		
otassium	к ₂ 0	1.42 ± 0.16	1.21 ± 0.12	1.07 <u>+</u> 0.01	1.09 <u>+</u> 0.11		
ron	Fe203	<u>7.9 ± 0.3</u>	<u>7.8_±1.1</u>	<u>6.5 ± 0.2</u>	<u>6.1 ± 0.3</u>		
		•	Average Concentration	a in Parts Per Million*			
bidium	Rb ₂ 0	72 <u>+</u> 9	66 <u>+</u> 2	64 <u>+</u> 13	60 ± 12		
sium	C3,0	4.2 <u>+</u> 0.8	3.8 <u>+</u> 0.7	4.8 ± 1.3	5.4 ± 0.8		
irium	BaO	<u>310 + 100</u>	270 + 80	580 + 190	<u>450 + 130</u>		
andium	se203	36 <u>+</u> 1	32 ± 4	31 <u>+</u> 3	27 <u>+</u> 4		
nthanum	La203	64 <u>+</u> 3	66 <u>+</u> 5	87 <u>+</u> 27	49 <u>+</u> 7		
erium	CeO2	<u>164 + 28</u>	$\frac{136 \pm 17}{2}$	<u>70 ± 4</u>	<u>48 ± 10</u>		
ropium	Eu 203	<u>2.9 ± 0.2</u>	<u>2.7 ± 0.3</u>	<u>1.9 + 0.1</u>	<u>1.6 ± 0.2</u>		
fniua	heo2	<u>6.6 ± 1.0</u>	<u>6.4 ± 0.9</u>	<u>3.1 ± 0.1</u>	<u>2.9 ± 0.5</u>		
norium	Th02	15.5 <u>+</u> 2.6	15.5 ± 0.5	14.1 <u>+</u> 0.6	11.2 ± 2.8		
ntalum	Ta205	3.8 ± 0.1	3.2 ± 0.4	2.6 <u>+</u> 0.3	1.3 ± 0.2		
romium	Cr203	<u>201 + 10</u>	<u>199 ± 33</u>	<u> 261 + 28</u>	<u>254 + 30</u>		
nganese	MnO	140 <u>+</u> 40	<u>256 ± 101</u>	230 + 92	173 ± 28		
balt	CoO	16 <u>+</u> 2	23 ± 5	<u>21 + 1</u>	25 ± 4		

COMPARISON OF PLEISTOCENE CLAYS WITH SHALES FROM THE SAME SOURCE REGIONS

* With group standard deviations.

TABLE II

		Kaolin Clays	from Nubia an	d the Suez Gulf	1	Nubia Sandstone Clays from the Aswan Region							
Epoch of De	position		Gretaceo	۳.		Lover Crétaceous							
Source Loca	ations	Kalabsha South of Aswan	Abu DaragAswanGulf of Suezdark greyClayclayT18T19T10		dark grey	Aswan light grey clay	<u>Specimens</u> top layer	from a Shaft near opening	Near Aswan 75 meters inside				
Specimen No	.	T13			Ť20	T 21	T22	T23					
					Concentrations of Oxides in P	ercent			<u> </u>				
Sodium	Na ₂ 0	0.16	0.38	0.76	0.10	0.11	0.58	0.42	0.08				
Potassium	к,0		0.15	0.10	1.19	1.15	1.00	1.08	1.13				
iren	Fe 203	0.74	1.02	1.17	6.3	2.5	7.5	7.3	4.8				
			Concentrations of Oxides in Parts Per Million										
Rubidium	Rb 20	23	- 1-21 2	*****	70	50	42	46	48				
Cesium	Cs 0				4.3	3.1	5.3	6.7	4.6				
Bariua	BaÖ	350		230	. =====		250	230	240				
Scandium	Sc,03	43	34	41	38	54	55	41	42				
Lanthanum	La203	128	92		85	85	68	65	69				
Cerium	CeO,	480	371	417	186	160	164	145	152				
Europiua	Eu203	4.5	3.1	2.7	2.8	3.4		2.8	2.4				
Rafnium	NEO,	21.2	15.9	15.7	7.0	8.0	6.1	7.6	7.9				
Thorium	Th0,	38.1	20.7	30.8	14.9	15.8	14.9	16.7	17.8				
Tantaluc	Ta203	15.4	10.5	12.0	3.8	5.5	1.9	2.2	2.6				
Chronium	Cr203	381	206	180	153	155	193	168	167				
Manganese	MnO	11	43	25	111	36	244	165	72				
Cobalt	CoO	6	3	5	33	6	22	17	14				

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TABLE III EGYPTIAN CLAYS DEPOSITED BEFORE FORMATION OF THE NILE

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TABLE IV

SHALES FROM OUTCHOPS WEAR THE WILE

ormation		Ma	ıdi. Parı	stica	•	Lean 1	ihélų.		Tekk I	A Shale				Theephal	te forma	tion		
poch of opewition		U	pper Eod	20¢		Upper Pa	lestens	atona Lover Faloocono		Companian to Heastrichtion								
ource Loca	tin	El F	éyu:z	Headi south of Ceiro			l-Sheehol Distri			aveine gyöt		Cebe Northea	L Alui Ka it of Qu	-		Coreyst Shars	<u>Gebe</u> waa Dia	<u>l Aveina</u> trict
pecimen W		TK17	TKLS	TE20	TS4	785	7513	TS14	753	T\$12	TAL	TAS	TA7	TA19	753	759	Ť\$10	T511
							C	oncontration	of Oxides is	Tercent		_						
odius	84 ₂ 0	1.76	1.67	1.67	2.43	2.78	2,24	2.25	1.57	3.50	8.44	2.79	3.44	1.59	2.91	4.39	4.14	4.41
otaasiwa	K.0	1.07	1.25	1.31	0.97	1.12	1.05	1.23	ŧ.75	1.07	1.17	6.94	0.99	2.96	1.16	1.42	1.36	1.19
ron	Yuju	5.24	8.62	6.39	5.92	5.82	6.45	6.1 9	5.20	4.77	5.08	4.56	5.48	8.06	9.17	7.04	7.26	5.70
							-	átrations of	Quides in Part	s Fer Million								
ubicium	Rb ₂ 0	6 8	66	65	72	43	65	58	57	48	51	53	49	43	65	73	55	50
esiun	(4 <u>2</u> 0	4.5	3.9	3.1	6.5	5.1	4.7	5.4	5.5	2.5	2.6	3.2	3.2	2.2	7.0	2.7	3.2	2.3
arlum	440	340	180	280	600-	520	300	390		199	600	230	350	*****	270	390	310	340
icandium	**203	35	32	28	33	23	26	27	37	24	23	30	29	32	33	28	37	21
.anthan um	La203	71	64	62	39	42	46	47	41	47	38	44	38	44	50	53	44	57
Cerium	CeO,	151	118	140	41	38	44	46	42	51	40	46	42	45	49	53	45	59
Luropius	2u203	2.9	2.8	2.3	1.9	1.4	1.6	1.6	1.3	1.6	2.4	Q.9	1.7	1.8	2.0	2.0	1.7	1.8
iafnius	8502	5.9	7.ū	5.4	3.5	2.3	2.8	3.0	3.7	3.0	2.3	2.5	ź.ż	3.1	2.9	4.7	3.9	6.0
Thorium	Tho, .	16.0	15.0	15.5	15.0	8.4	10.3	11.0	13.2	11.3	8.4	11.9	11.8	8.7	11.6	12.7	12.0	12.9
Tuntalus	Ta.,0,	3.7	2.9	3.0	1.5	1.4	1.1	1.3	1.4	9.9	0.8	1.1	0.8	1.0	1.4	1.3	1.1	1.1
Chronius	Cr30,	221	215	161	290	217	255	255	222	190	245	227	333	177	329	396	402	193
Manganese	Sta)	151	353	264	133	177	192	192	48	33	75	49	47	- 354	177	225	164	50
			29				29	24								-	-	10

TABLE V

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SHALES FROM GEBEL ABU TARTOUR AT THE KHARGA OASIS

Formation		Da'	khla Sha	1e	Phosphate Formation		Variegated Shales					
Epoch of Deposition		Lower Paleocene			Crétaceous	Upper Cretaceous (Cenomenian)						
Specimen No	.	TK13	TK14	TK15	TK10 ·	TK3	TK5	TK6	ŤK 7	TK8		
	<u></u>				Concentrations of Oxides in Percent					نديج ميتينيسيوني		
Sodium	Na ₉ 0	0.10	0.18	0.14	0.06	0.14	0.13	0.09	0.13	0.09		
Potassium	к,0	1.09	1.61	1.54	0.89	3.27	3.72	1.96	4.04	2.37		
Iron	Fe203	5.58	6.09	8.69	4.83	9.09	9.05	8.52	9.31	7.99		
					Concentrations of Oxides in Parts Per Hillion							
Rubidium	Rb ,O	43	46	52	28	95	76	62	62	66		
Cesium	Cs ₂ 0	2.2	1.8	2.5	1.5	1.5	1.7	1.3	1.5	1.6		
Barium	5a0		46Ô	~~*****	330		190	520	220	400		
Scandium	Sc 203	26	22	35	31	29	32	20	30	31		
Lanthanum	ta203	46	52	61	52	44	50	77	48	58		
Cerium	Ce02	87	100	144	109 .	72	91	172	91	99		
Europium	Eu203	1.8	1.8	1.7	1.3	1.6	2.2	1.3	2.0	2.3		
Hafnium	HEO2	2.4	12.6	4.2	3.1	4.3	5.6	15.8	5.0	5.6		
Thorium	Th02	12.4	15.6	15.5	11.5	9.6	12.1	16.7	10.2	16.8		
Tantalum	Ta205	1.4	1.5	1.8	1.1	1.5	2.4	2.0		0.7		
Chromium	Cr203	192	196	245	195	139	157	113	149	188		
Manganèse	Mn0	215	428	96	145	330	384	253	267	176		
Cobalt	CoO	21	20	15	22	23	24	17	22	18		

TABLE VI

ANALYSES O	? Some	ANCIENT	EGYPTIAN	POTTERY
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		C	compositio	nal Group	1	Compositiona	1] Group 2	Compositio	onal Grou	ıp 3
Source Loc	ation		Sakkara	/	lexandria	Alexan	dria	A1(exandria	
Period		Archaic	Pyramid	26th	Early					<u></u>
		Period	Period	Dynasty	Islamic	Ptolemaic	Coptic	P	colemaic	
Specimen N	la.	P1	P9	89	P2	23	P5	P4	P6	P7
<u> </u>		*****				Compositions of Oxider	in Percent			
Sodium	Na ₂ 0	2.28	2.25	2.60	1.80	1.75	1.25	1.21	1.34	0.63
Potassium	к , о	1.80	2.18	1.50	1.50	1.92	2.01	2.37	2.71	1.27
l ron	Fe203	10.20	8.50	11.00	11.60	6.55	5.96	7.83	7.12	7.05
						Compositions of Oxides in 1	Parts Per Hillion			
Rubidium	Rb o	70	57	60		42	62	124	130	61
Cesium	Cs ₂ 0	1.2	1.2	1.4	0.9	1.0	1.4	8.2	5.7	3.5
Barium	840	410	640	490	600	450	400	410	370	220
Scandium	se203	37	33	40	42	25	21	34	31	52
Lanthanum	La ₂ 03	46	41	42	44	49	38	39	39	19
Cerium	CeO,	97	83	92	92	92	76	\$1	77	47
Europium	Eu203	2.5	2.6	3.2	3.0	1.7	1.7	1.5	1.5	0.7
Hafnium	HEO,	6.7	6.8	6.8	7.5	9.1	5.4	4.2	4.2	1.8
Thorium	Th02	9.2	7.1	7.8	8.0	8.8	7.4	12.3	11.6	6.6
Tantalum	Ta205	2.6	2.7	2.4	3.1	3.3	2.5	1.5	1.5	0.9
Chromium	Cr203	236	184	249	281	169	126	520	350	330
Manganese		2050	1180	1760	1880	950	960	1140	1510	890
Cobalt	CoO	46	38	48	54	31	26	40	39	37

TABLE VII

COMPARISON OF THE AVERAGE OXIDE CONCENTRATIONS IN THE NILE ALLUVIUM SPECIMENS AND IN FOUR ANCIENT POTSHERDS* OF SIMILAR COMPOSITION

		Nile Al Speci			atching nt Pottery
		Average	Concentration	s of Oxides i	n Percent
Sodium	Na ₂ 0	1.42 <u>+</u>	0.18	2.2	3 ± 0.33
Potassium	к ₂ 0	1.32 <u>+</u>	0.13	1.7	2 <u>+</u> 0.32
Iron	Fe203	11.2 <u>+</u>	0.3	10.	3 <u>+</u> 1.4
		Average Conce	ntrations of (xides in Par	te Per Millio
Rubidium	^{ℝЬ} 2 ⁰	50 <u>+</u>	15	62	± 7
Cesium	Cs20	1.3 <u>+</u>	0.4	1.2	± 0.2
Barium	BaO	510 <u>+</u> :	30	540	<u> </u>
Scandium	Sc203	40 <u>+</u> :	1	38	<u>+</u> 4
Lanthanum	La203	43 <u>+</u> 3	10	43	<u>+</u> 2
Cerium	Ce02	108 <u>+</u> :	33	91	<u>±</u> 6
Europium	^{Eu} 2 ⁰ 3	2.9 <u>+</u> (0.2	2.8	<u>+</u> 0.3
Hafnium	Hf02	7.3 <u>+</u> (0.7	7.0	<u>±</u> 0.4
Thorium	^{Th0} 2	6.7 <u>+</u> (0.7	8.0	<u>+</u> 0.9
Cantalum	^{Ta} 2 ⁰ 5	2.8 ± (0.8	2.7	<u>+</u> 0.3
Chromium	^{Cr} 2 ⁰ 3	220 <u>+</u> 2	22	238	<u>+</u> 40
langanese	MnO	1660 <u>+</u> 1	L30	1720	<u>+</u> 380
Cobalt	Co0	46 <u>*</u> 1	L	46	<u>+</u> 7

* Pottery Compositional Group 1.

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Figure 1. Geological Map of the Nile Valley in Upper Egypt (Reproduced from <u>The Geology of Egypt</u> by R. Said, with the kind permission of Dr. Said.)

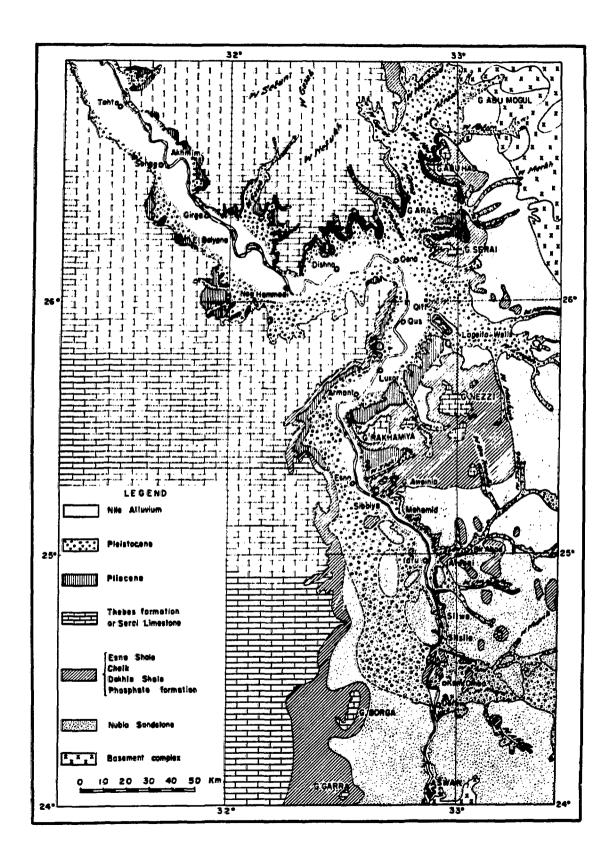
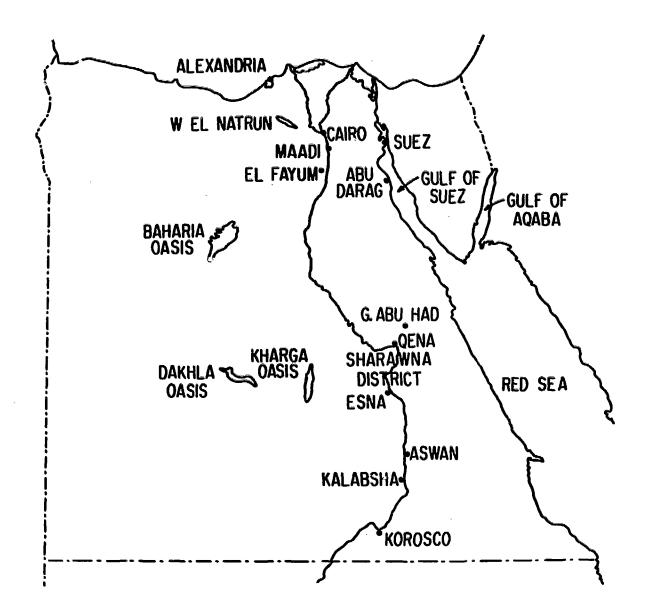
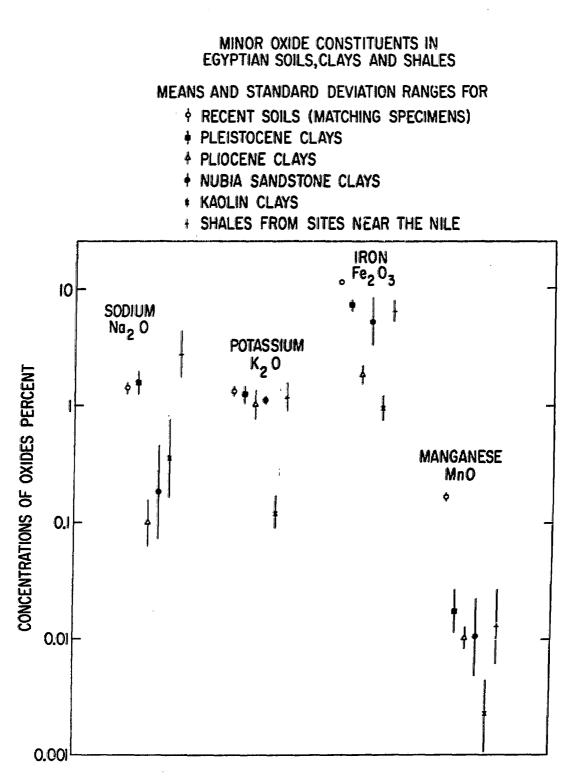


Figure 2. Sample Source Locations in Egypt



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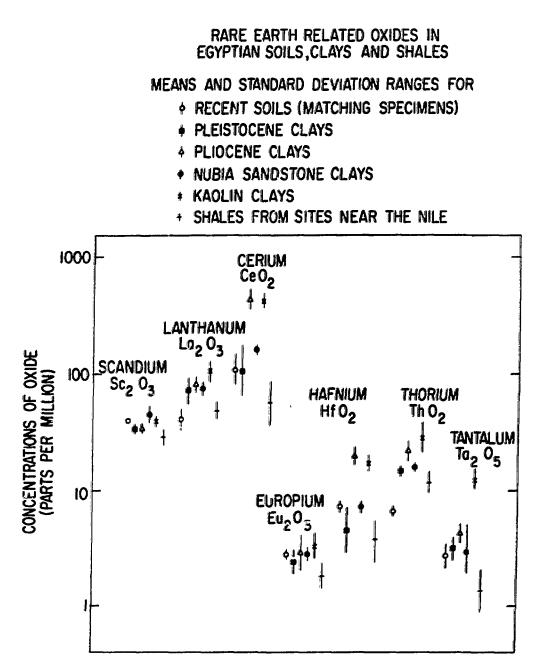


Figure 4