

## Deformation due to the 17 August 1999 Izmit, Turkey, earthquake measured from SPOT images

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[1] The geometry of the ruptured areas and the coseismic slip distribution data are key to highlighting the behavior of seismic faults. This information is generally retrieved from field investigations and geodetic measurements or synthetic aperture radar (SAR) interferometry. Here we show that SPOT images can also be used to accurately map the fault zone and to determine the slip distribution by subpixel correlation of images acquired before and after an earthquake. The measured slip includes the contribution of possible distributed shear that might not be clearly expressed in surface ruptures and smoothes out possible along-strike variability due to near-surface fault complexities. We apply the technique to the  $M_s = 7.4$ , 1999, Izmit earthquake. Our results reveal a <100-m-wide and very linear fault zone that can be traced for 70 km from Gölcük to Akyazi, along which supershear rupture has been inferred. The obtained slip distribution compares well with the field measurements and is consistent with ground deformation measured at some distance from the fault zone using SAR images. Very little deformation was accommodated off the main fault plane. Maximum slip is observed near Sapanca lake at a small fault jog that has probably influenced rupture propagation. *INDEX TERMS*: 1242 Geodesy and Gravity: Seismic deformations (7205); 1640 Global Change: Remote sensing; 5475 Planetology: Solid Surface Planets: Tectonics (8149); *KEYWORDS*: earthquake, deformation, Izmit, offset

### 1. Introduction

[2] Fault ruptures geometry and coseismic displacements data are key to highlighting the behavior of seismic faults. This behavior is of interest in earthquake source studies because fault geometric complexities are known to influence rupture dynamics [Sibson, 1986; Harris *et al.*, 1991; Barka and Kadinsky-Cade, 1988; Bouchon and Streiff, 1997; Aochi *et al.*, 2000; Harris and Day, 1999; Kase and Kuge, 1998; Harris *et al.*, 2000], hence seismic radiation, and because the coseismic slip distribution can be used to test dynamic rupture models [Bouchon and Streiff, 1997; Aochi *et al.*, 2000]. A preliminary attempt to model spontaneous rupture during the Izmit earthquake taking into account the fault segmentation has been performed by Harris *et al.* [2000]. Conversely, such data are necessary to constrain the determination of the source kinematics from seismological records because there are various trade-offs between the geometric characteristics of the fault, the slip, and the rupture time distributions [Hernandez *et al.*, 1999]. A good knowledge of fault geometry and coseismic slip is also needed to estimate coseismic stress changes and investigate aftershock triggering mechanisms and implications for the probability of rupture of other nearby faults [Stein, 1999; Hubert-Ferrari *et al.*, 2000; Parsons *et al.*, 2000].

[3] Geometry of fault ruptures and coseismic slip is generally obtained from field investigations and geodetic measurements. Field investigations consist of thorough mapping of surface breaks and measuring coseismic slip from dislocated features that can be matched across these breaks [Yeats *et al.*, 1997; Weldon *et al.*, 1996]. However, because of near-surface fault complexities the determination of the total slip across the fault zone is not straightforward. In addition, the zone of coseismic strain might be significantly broader than the zone where ruptures are clearly seen at the surface. In the case of the Landers earthquake it has been observed that anelastic shear may be distributed in a zone possibly as broad as

1–2 km and may amount to as much as 20% of the total slip across the fault zone [Yeats *et al.*, 1997; McGill and Rubin, 1999]. It is therefore not straightforward to assess whether the slip variability measured at the surface only relates to these near-surface complications or may reflect slip heterogeneity at depth.

[4] For these various reasons, any technique capable of providing reliable measurements of ground displacements near the fault zone would be most useful. Geodetic data are generally too sparse to constrain tightly the slip distribution. Differential synthetic aperture radar (SAR) interferometry [Massonnet *et al.*, 1993; Zebker *et al.*, 1994] generally fails in the near-fault zone where large displacements result in poorly correlated images and eventually signal decorrelation [Michel *et al.*, 1999]. The use of subpixel correlation of SAR amplitude images can partially supplement interferometry as it provides an unambiguous measurement of both the track-parallel and track-perpendicular components of the ground displacement [Michel *et al.*, 1999; Peltzer *et al.*, 1999]. Unfortunately, temporal decorrelation of ERS amplitude images available for the Izmit earthquake is too large, and SAR images do not yield significant measurements near the faults. Following early attempts [Crippen, 1992], we have explored the possible use of optical imagery to measure ground displacements and developed specific algorithms [Van Puybroeck *et al.*, 2000]. This technique is applied here as a complement to radar interferometry for analyzing ground displacement during the Izmit earthquake.

### 2. Setting of the Izmit Earthquake

[5] The Izmit earthquake ruptured a segment of the North Anatolian fault (NAF) at the eastern tip of the Marmara Sea (Figure 1). This fault segment is part of a northern branch of the NAF which is thought to have accommodated ~70 km of slip since ~5 Myr [Armijo *et al.*, 1999]. GPS measurements [Straub and Kahle, 1997; McClusky *et al.*, 2000] anterior to the Izmit earthquake revealed interseismic straining corresponding to  $27 \pm 1$  mm/yr slip rate and a fault locking depth of ~20 km [Reilinger

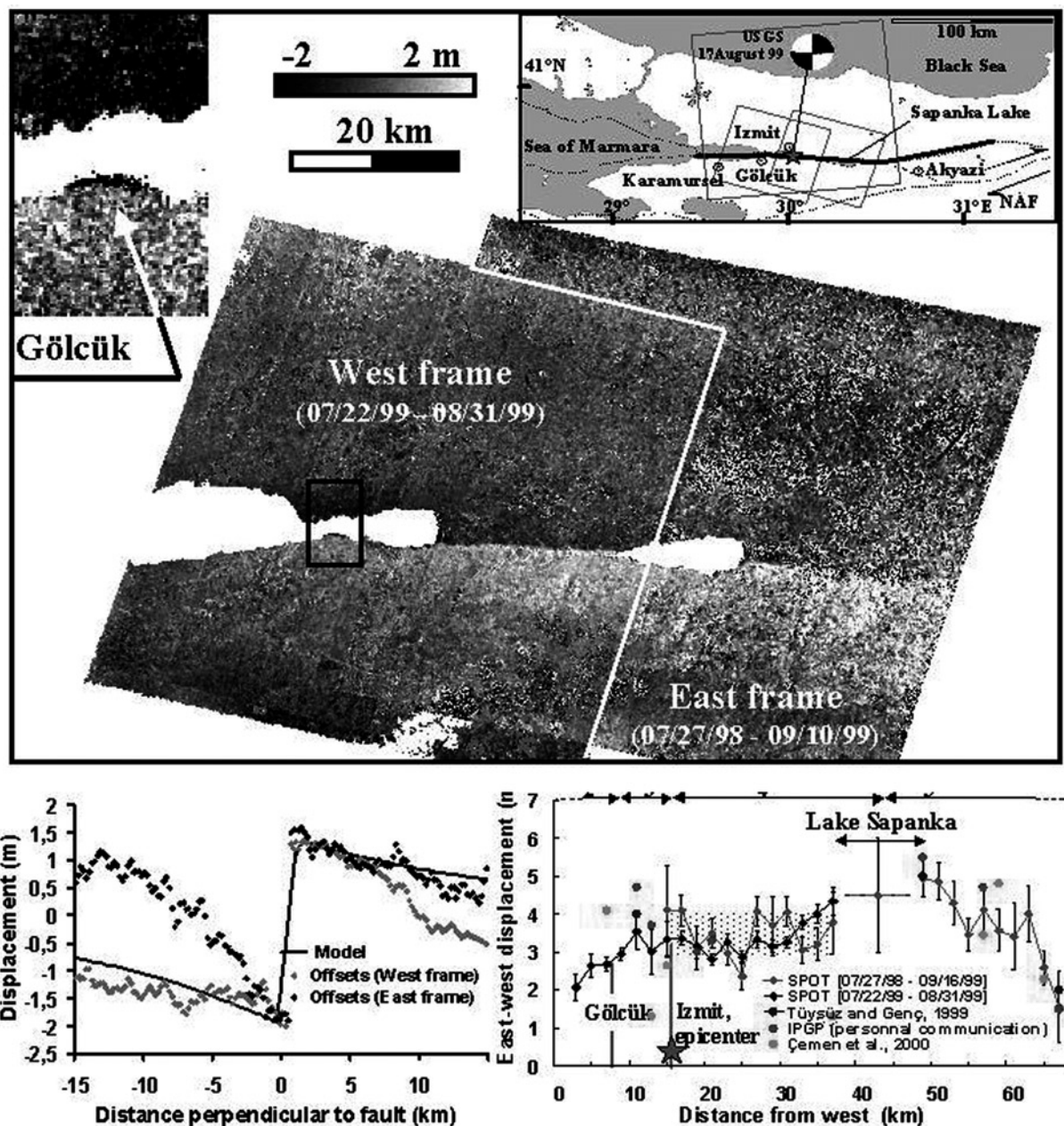


Figure 1. Ground displacement mapped from SPOT images. (a) E-W component of the displacement field computed from image pairs (KJ 102/268, SPOT-1 of 22 July 1999 and SPOT-2 of 31 August 1999) and (KJ 103/268, SPOT-2 of 27 July 1998 and SPOT-2 of 16 September 1999). The very linear discontinuity between Gölçük and Akyazi coincides with the ruptures described from field survey. No deformation is visible out of this main fault zone. N-S offsets are below the noise level. Tectonic setting, with main active faults, of the  $M_s = 7.4$ , 17 August 1999 Izmit earthquake is given in inset. Focal mechanism is from USGS catalog, and epicenter (star) from Kandilli Observatory. Shaded boxes denote location of SPOT and ERS images. Black lines show approximate location of ruptures described from field investigations. (b) Stack of all profiles across fault zone from Izmit to Sapanca lake. Offsets show a clear jump at the fault trace, indicating a fault zone  $<200$  m wide. Mean slip is  $3.5 \pm 0.2$  m. The two profiles differ due to uncertainty on SPOT changing attitude. Shaded line shows theoretical displacements derived from elastic modeling (Table 1). (c) Slip distribution between Gölçük and Akyazi measured from 1-km-wide swaths across the fault. Uncertainties are estimated from the variability of offsets on each side of the fault. The estimated slip at Sapanca lake is ascribed a large uncertainty because it is based on linear extrapolation of offsets across the lake. Field measurements are shown for comparison. See color version of this figure at back of this issue.

*et al.*, 2000]. The epicenter of the  $M_w = 7.4$ , 17 August earthquake is located close to Izmit [Toksök *et al.*, 1999; Barka, 1999] (Figure 1). The source mechanism is purely strike-slip faulting on a vertical E-W fault [Barka, 1999; Tüysüz and Genç, 1999; Çemen *et al.*, 2000]. According to local strong motions and teleseismic

records the hypocentral depth was  $\sim 15$  km [Çemen *et al.*, 2000; Ellsworth and Celebi, 1999]. The total seismic moment is estimated between 1.0 and 2.4  $10^{20}$  N m [Toksök *et al.*, 1999; Çemen *et al.*, 2000]. Surface ruptures were observed for  $\sim 120$  km extending east-west from Karamürsel to near Düzce [Barka, 1999; Tüysüz and



**Table 1.** Fault Model Derived From SPOT Offsets and ERS-1 SAR Interferometry<sup>a</sup>

Fault Segment	Center Latitude, °N	Center Longitude, °E	Length, km	Azimuth, °E	Slip, cm	Averaged Slip, <sup>b</sup> cm	Depth, km
1, Marmara	40.719	29.462	16.64	-92.203	200	100	15
2, Karamürsel-Gölcük	40.710	29.682	22.11	-93.318	315	300	11.0
3, Gölcük-Izmit	40.718	29.869	10.00	-83.56	315	300	13.6
4, Izmit-Sapanca	40.716	30.091	27.86	-87.690	349	330	8.9
5, Sapanca-Akyazi	40.701	30.403	25.08	-81.199	220	250	8.8
6, Akyazi-Duzce	40.720	30.739	32.61	-101.310	100	130	10

<sup>a</sup>All fault planes are vertical and reach the surface. Angles are given in degrees, and azimuth is given counterclockwise from due north.

<sup>b</sup>Averaged value from *Reilinger et al.* [2000]; see text for discussion.

*Genc, 1999; Çemen et al., 2000*]. The ruptures probably extend west of Karamürsel, in the Gulf of Izmit and Marmara Sea as suggested from the aftershocks [*Parsons et al., 2000*]. From Karamürsel to Akyazi the fault strikes approximately E-W azimuths, is relatively linear with small jogs, and has accommodated right-lateral slip between 1 and 5 m [*Barka, 1999; Çemen et al., 2000; Barka et al., 2000*]. This fault portion is probably responsible for the main subevent in the bimodal source function [*Çemen et al., 2000; Ellsworth and Celebi, 1999*] and is characterized by a supershear rupture velocity [*Bouchon et al., 2000*]. It has been proposed that the bimodal source would also be reflected in the surface slip distribution with maximum slip east of Sapanca and near Golcük [*Çemen et al., 2000*]. The rupture of the easternmost segment, which strikes about N65°E east of Akyazi, was apparently triggered late in the process [*Barka, 1999*]. Field measurements indicate <1.5 m of right-lateral slip along this segment [*Barka, 1999; Tüysüz and Genc, 1999; Barka et al., 2000*].

### 3. Determination of the Fault Geometry and Slip From SPOT Images

#### 3.1. Method

[6] Ground deformation during an earthquake may be measured from comparison of optical images acquired before and after the seismic event [*Van Puymbroeck et al., 2000*]. We use pairs of SPOT images with near-vertical incidence in order to minimize stereoscopic disparities. The two images are first resampled in the same cartographic projection using a digital elevation model (DEM), a physical model of image acquisition, and an accurate interpolator. In our study we use a DEM derived from ERS SAR interferometry. During that initial processing, some artifacts due to the SPOT push-broom system are corrected together with misregistrations due to the changing attitude of SPOT during the acquisition of each image. Residual offsets are next computed from the phase shift of the Fourier transform of a sliding window. It yields the two horizontal components of the ground displacement and an estimate of errors. Spatial sampling and the accuracy of measurements mainly depend on the local level of correlation of the images. For the Izmit earthquake we get independent measurements every 160 m and an average uncertainty of 0.25 m (1/40th of the pixel size). Offsets also include residual artifacts due to errors and undersampling of the roll, pitch, and yaw characterizing SPOT attitude. This latter source of error may introduce a significant apparent warping at spatial wavelengths larger than a few kilometers [*Van Puymbroeck et al., 2000*].

#### 3.2. Application to the Izmit Earthquake

[7] We have focused our study on the main fault strand between Karamürsel and Akyazi. We used two pairs of SPOT images, one pair of ERS SAR images and a DEM derived from ERS images (Figure 1). Each pair spans the coseismic phase together with 15–25 days of postseismic deformation. The

images were first corrected to account for various artifacts that cause misregistrations unrelated to ground deformations. The remaining offsets were then measured from the subpixel correlation technique described above. The E-W component (Figure 1a) shows a prominent approximately E-W discontinuity between Izmit and Akyazi that closely follows the surface ruptures mapped in the field [*Tüysüz and Genc, 1999; Çemen et al., 2000*]. Even the surface ruptures along the coast line from Gölcük to Izmit are visible (inset in Figure 1a). The N-S component of the displacement field is everywhere below the noise level, as expected from a nearly pure strike-slip motion. In order to investigate ground deformation in the fault zone we have drawn profiles perpendicular to the fault strike, by stacking all data along 1-km-wide swaths. All the profiles show a clear jump at the fault trace. Figure 1b shows for example a stack of the 26 profiles along the Izmit-Sapanca fault segment, retrieved from the two pairs of SPOT images. The two sets of data show the same jump at the fault trace, but the patterns are actually quite different. The difference results from uncertainties on the attitude of SPOT during image acquisition. Fortunately, the satellite attitude does not vary rapidly enough to introduce as sharp gradients as those induced by coseismic deformation. The determination of the fault slip at the surface rupture is not biased. SPOT offsets suggest that the fault zone is nowhere wider than ~200 m, with a clear right-lateral offset of 3.5 m on average. This value is consistent with the 3–4 m slip reported from field investigations along this fault segment [*Tüysüz and Genc, 1999*]. The value is also in agreement with the displacement measurements by *Çemen et al.* [2000], who reported a gradual decrease in displacements from ~4 m near Golcük to 1.3 m ~5 km west of Sapanca lake. They also reported that displacements increases again eastward and reaches to ~5 m east of the town of Arifiye. Figure 1c shows the slip pattern derived from the 1-km spaced profiles drawn in the two offset fields. Average uncertainties of the measurements are 0.25 and 0.70 m for the west and east frames, respectively. Comparison between the measurements made from the two offsets fields suggests that measured uncertainties are probably slightly underestimated. The field measurements in Figure 1c appear to be close to our measurement of the total offset across the fault zone. Our estimate of the slip between Gölcük and Izmit seems small compared to the field measurements. This is probably due to uncertainty associated with the interpolation of offsets across Izmit Bay. The general consistency between field measurements and those made from SPOT images and the fact that the fault zone seems everywhere narrower than one pixel in the offset field (160 m) indicate that coseismic fault slip was entirely accommodated along the ruptures seen at the surface. We can conclude that there was little if any coseismic deformation taken up by distributed shear off the main fault's trace. Because field measurements generally show a complex fault geometry and a high variability of slip, with slip possibly varying by as much as a few meters over a few hundreds meters [*Yeats et al., 1997; Weldon et al., 1996; McGill and Rubin, 1999*], we should not expect a perfect match between 1-km averaged values of the

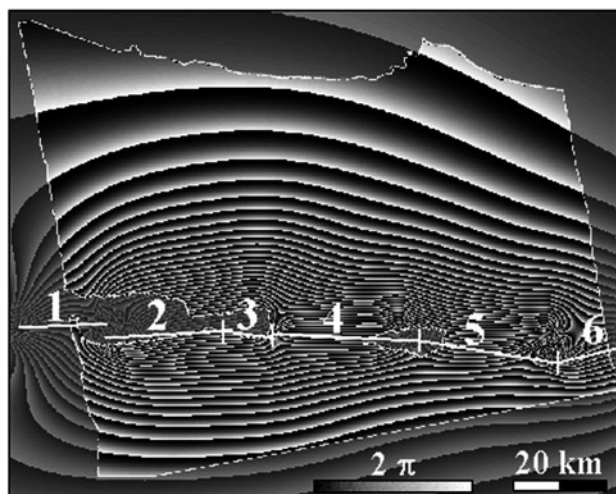
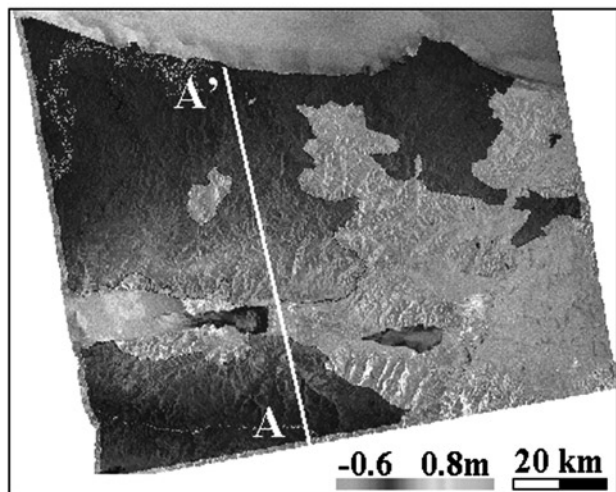
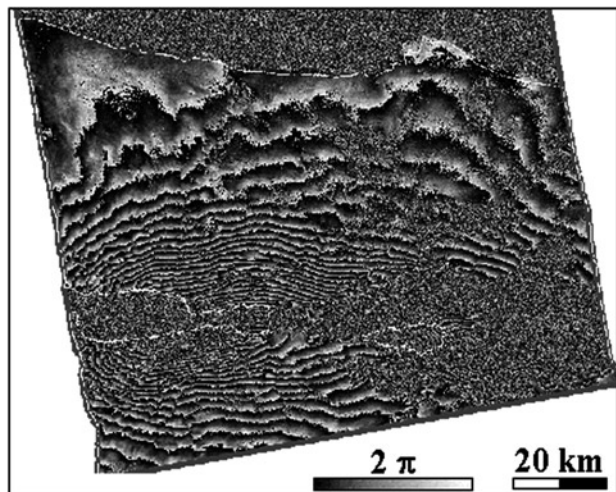
displacement across the whole fault zone obtained from our technique and local field measurements. In that respect, the surface ruptures associated with the Izmit earthquake seem remarkably simple. Note also that the generally high slip variability at the small scale implies that the estimate of the mean slip highly depends on the sampling in the field, so that a reliable estimate requires a very detailed survey [Yeats *et al.*, 1997; Weldon *et al.*, 1996]. Our technique is thus best adapted to yield a reliable estimate of the mean slip, together with some informa-

tion about possible distributed shear in the fault zone not clearly expressed in surface ruptures. It is also noteworthy that the slip variability measured from offsets (Figure 1) is much more regular than that derived from GPS data by Reilinger *et al.* [2000] though the values averaged at the segment scale are in good agreement (Table 1). This suggests that the model of Reilinger *et al.* [2000] is somewhat overdetermined as the number of free parameters in their model is greater than the number of GPS measurements available.

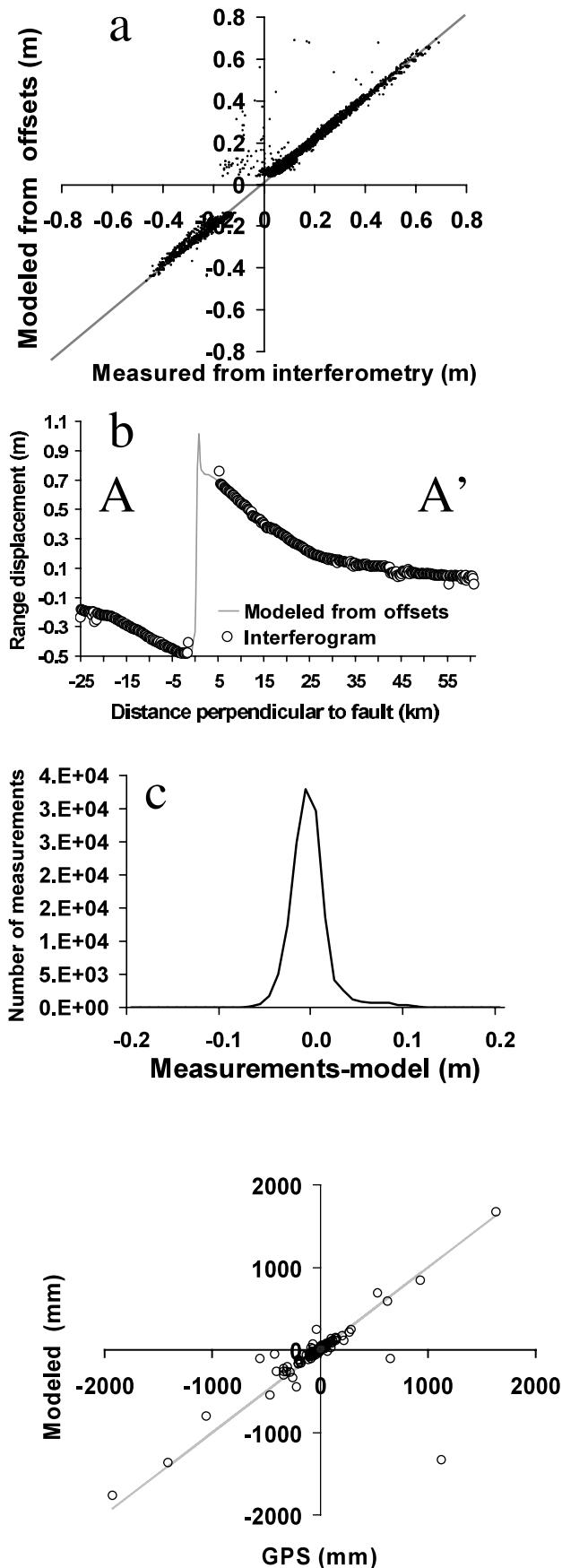
[8] In the present case we find mean slip of  $3.2 \pm 0.5$  m along the Karamürsel-Gölcük segment,  $3.4 \pm 0.4$  m along the Gölcük-Izmit segment,  $3.5 \pm 0.2$  m along the Izmit-Sapanca segment, and  $2.2 \pm 0.6$  m along the Sapanca-Akyazi segment (Table 1). The slip distribution obtained from our technique is rather smooth, with the zone of maximum slip,  $\sim 5$  m, being observed around Sapanca lake. Unexpectedly, it seems that the near-surface fault slip tends to be maximum at the junction between the fault segments. The slip tapers gradually from Sapanca to Akyazi. Farther east, it continues to trail off from Akyazi to Duzce along N56°E striking surface ruptures with slip between 1 and 1.5 m [Barka, 1999; Tüysüz and Genc, 1999]. We have not tried to document this zone of low slip from the SPOT images. It should be noted that field investigations around Sapanca lake have revealed much smaller offsets than those measured from SPOT images. This may suggest that near the fault jog at Sapanca lake, part of the coseismic deformation was accommodated in a distributed shear zone so that only a fraction of the total slip was accommodated in the main fault trace. More generally, the smoothness of the fault slip measured from the SPOT images suggests that the irregular fault slip distribution measured in the field reflects near-surface complexities in a shear zone only a few hundred meters wide.

#### 4. Consistency with SAR Measurements

[9] We have complemented our study using ERS-1 SAR images by computing SAR interferograms [Massonnet *et al.*, 1993; Zebker *et al.*, 1994; Michel *et al.*, 1999]. This technique provides the near-vertical component of the ground displacement (Figure 2). SAR interferograms are generally very noisy near the fault but provide



**Figure 2.** (opposite) Measured and modeled interferograms. (a) Differential SAR interferogram computed from ERS-2 SAR images (orbit frames 22556/815 and 23057/815) acquired on 13 August 1999 and 17 September 1999, respectively. A fringe represents 2.8 cm of ground shift away from satellite along the radar line of sight. Areas with a high level of noise result from temporal decorrelation. Near-fault areas are also poorly coherent due to either ground shaking or signal saturation. Interferogram is also affected by low-frequency noise (for wavelength larger than  $\sim 10$  km) with an amplitude of a few fringes induced by atmospheric artifacts and residual errors in determination of orbits [e.g., Reilinger *et al.*, 2000]. Orbital parameters tuned to remove a residual plane that cannot be explained by coseismic ground deformation. This linear correction amounts to three fringes in azimuth and four fringes in range. (b) Unwrapped interferogram. Phase unwrapping removes the  $2\pi$  ambiguity on measurements and yields relative displacements. Because the fault zone cuts across the interferogram, areas north and south of the fault zone are not related to the same reference frame, they were adjusted in order to fit elastic modeling. (c) Synthetic interferogram computed from the segmented fault model given in Table 1. We use an elastic half-space dislocation model and consider vertical faults that reach the surface. Geometry and mean slip along the Karamürsel-Gölcük, Gölcük-Izmit, Izmit-Sapanca, and Sapanca-Akyazi segments were constrained from the offsets. The fault extents at depth and dip angle were adjusted for each segment. See color version of this figure at back of this issue.



valuable constraints on ground deformation at distance from the fault trace [Michel *et al.*, 1999]. Our measurements are in good agreement with the interferograms of Wright *et al.* [2001] and Reilinger *et al.* [2000] as they differ by less than two fringes. The discrepancy is below the noise. It may result from differences in the baseline tuning procedure. About 60% of the interferogram could be unwrapped satisfactorily, but the displacement closer than  $\sim 10$  km from the fault zone could not be resolved (Figure 2). We cross checked the two sets of data and tried to derive some simple fault models. We used dislocation modeling [Okada, 1985], which considers rectangular fault planes embedded in an elastic half-space. The fault strand between Karamürsel and Akyazi is described with four fault segments, which were forced to follow the fault trace evidenced from the SPOT images. In order to adjust the gradient of fault-parallel displacements seen from the SAR interferogram the faults must be dipping steeply (between  $80^\circ$  and  $100^\circ$ ). In the forward modeling, we have assumed coseismic slip on each fault segment equal to the average slip obtained from the SPOT images (Table 1). We have introduced the Marmara and Akyazi-Duzce fault segments as proposed by other investigators [Barka, 1999] and could not refine their description from our data. The fault description was further improved using the SAR interferogram. We adjusted the slip and the depth the modeled fault by minimizing, using a least squares criterion, the distance between the modeled and measured displacement in areas where the interferogram could be unwrapped. In addition, the fringe pattern was adjusted visually in the areas where some fringes can be seen, although they could not be unwrapped. The SAR interferogram put constraints on the extent of these faults at depth and on the faults that extend within the Marmara Sea west of Karamürsel (segment 1 in Table 1). The characteristics of the segment east of Sapanca lake are not that well constrained because of the noise on the interferogram. The model appears to meet the displacements measured from the SAR interferogram to better than 5 cm (RMS) (Figure 3). The proposed model is also in agreement with the 51 coseismic GPS measurements available from Reilinger *et al.* [2000] (Figure 3d). This result validates the slip distribution and fault geometry derived from the SPOT images. It moreover suggests that this simple fault model is sufficient to explain the SAR data. It does not mean that the fault geometry and slip distribution are as simple as assumed in the model but that the SAR data do not require a higher degree of complexity. In particular, the interferogram can tell us about the dip and extent at depth of the ruptured area but are not capable of revealing a more complex slip heterogeneity at depth than that assumed in the model. The inferred depth of faulting of 12–14 km for the segments west of Izmit and  $\sim 9$  km east are consistent with the hypocentral depth of the main shock and major aftershocks [Toksök *et al.*, 1999]. The model in Table 1 corre-

**Figure 3.** (opposite) Comparison of ground displacement measured from interferometry and computed from the fault model of Table 1. (a) The comparison is made along profile AA'; See Figure 2b for location of profile. (b) Plot of observed versus theoretical displacements for the whole area covered by the unwrapped interferogram. Values closely follow the 1:1 linear trend and meet the origin. It attests to the quality of the adjustment and validates the assumption of no displacement in the far field. The points with negative displacements correspond to the area south of the fault. The reference frame for this area is adjusted so that these points also fall on the linear 1-1 trend trough the origin. (c) Histogram of the difference between the observed and modeled displacements. It shows that the model fits the measurements to better than 5 cm. Nearly all the information contained in the interferogram is accounted for by the simple six-segment fault model of Table 1. (d) Comparison between the proposed model and 51 coseismic GPS measurement available [Reilinger *et al.*, 2000]. RMS error (6.5 cm) is in agreement with noise on measurements.



sponds to a static moment of about  $1.1 \times 10^{20}$  N/m<sup>2</sup> in the lower range of seismic estimates.

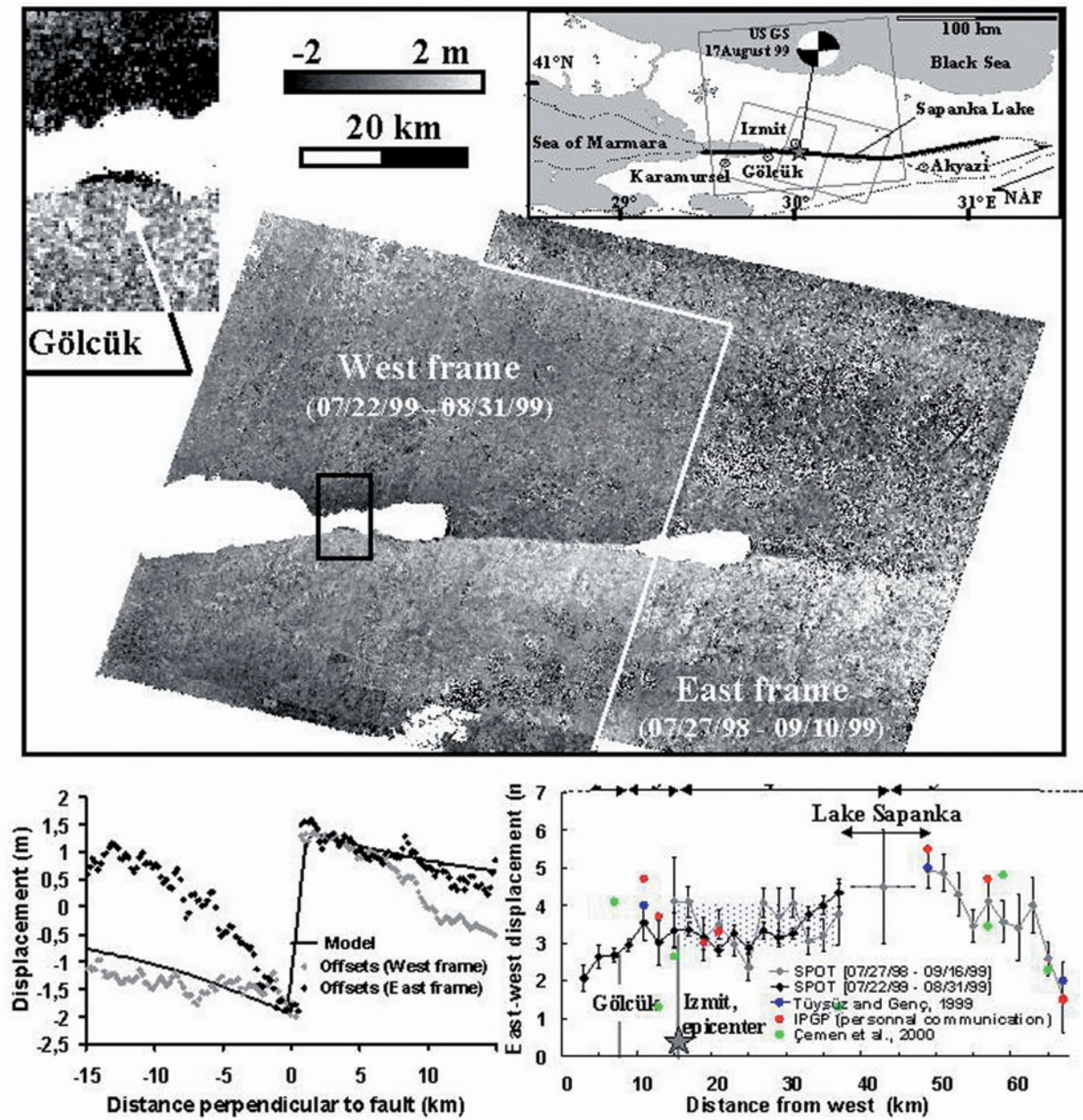
## 5. Implications and Conclusion

[10] Our study demonstrates that the geometric disparities between SPOT images (offsets) acquired before and after an earthquake provide constraints on the geometry of the fault zone, with an accuracy of the order of 100 m, and on the average fault slip, with an accuracy of  $\sim 50$  cm for 1-km spaced measurements. It allows some averaging of high-frequency slip variations induced by the near-surface fault complexity and yields a measurement of the total slip accommodated across the whole fault zone, including possible off-the-main-fault deformation. Applied to the Izmit earthquake, our technique reveals a higher coseismic slip at the fault junction near Sapanca lake and possibly also near Izmit and Gölcük. Our data thus tend to support the bimodal slip distribution inferred from the field measurements [Çemen *et al.*, 2000], although we observe a much smaller slip variation. These junctions are associated with small changes in the fault strike ( $3^\circ$ – $4^\circ$ ) and small extensional jogs. These minor geometric complexities may have influenced the dynamic rupture, possibly through some control on the preearthquake shear stress distribution or shear strength. The very linear geometry of the Izmit-Sapanca and Sapanca-Akyazi segments may have favored supershear rupture velocity [Bouchon *et al.*, 2000]. This technique should not be seen as a substitute to field investigations since it cannot provide information about the complexities of the fault geometry and the slip distribution at a scale smaller than a few tens of pixels.

[11] At the moment, the SPOT technique only provides information about the location of surface ruptures and slip across the fault trace because the offset field is biased by the satellite motion which is not known with enough accuracy. It does complement but not substitute field measurements. A finer description of the satellite attitude than that provided for SPOT images might allow us to overcome this limitation. The technique might then become even more powerful and could reveal the fault's geometry at depth which remains very poorly constrained from available techniques (GPS and SAR in particular). The technique may also gain further interest with the advent of high-resolution satellite images and might also be applied to archived airborne photography.

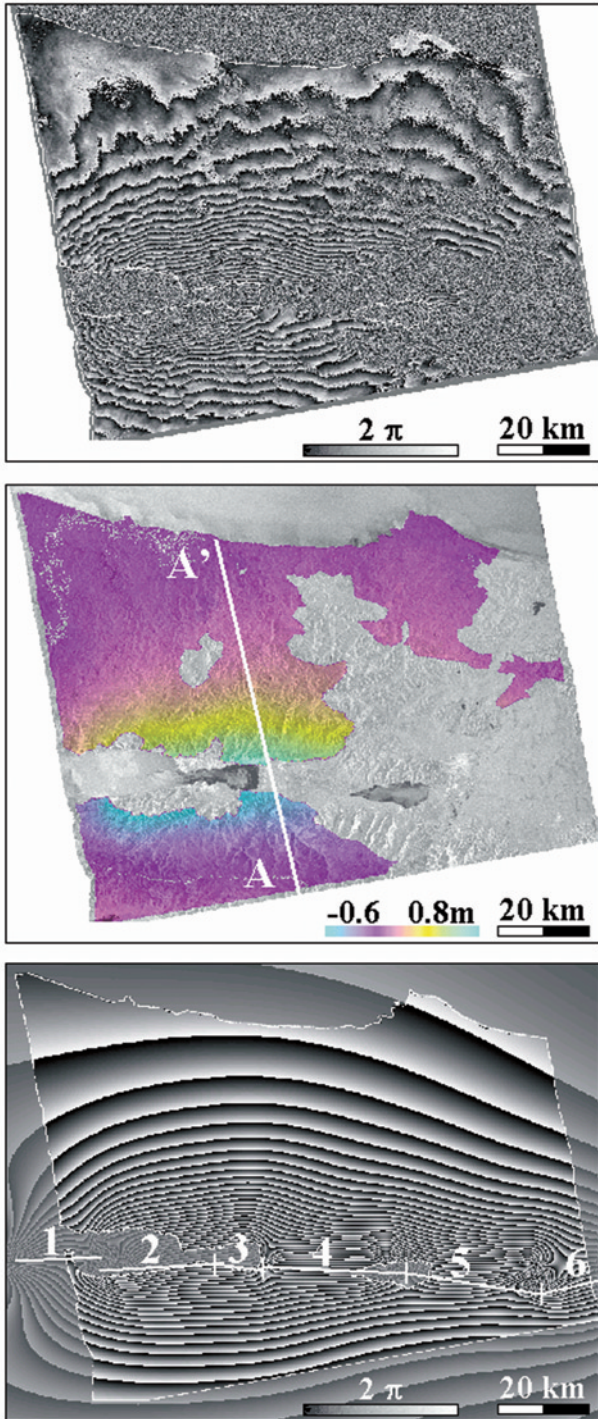
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**Figure 1.** Ground displacement mapped from SPOT images. (a) E-W component of the displacement field computed from image pairs (KJ 102/268, SPOT-1 of 22 July 1999 and SPOT-2 of 31 August 1999) and (KJ 103/268, SPOT-2 of 27 July 1998 and SPOT-2 of 16 September 1999). The very linear discontinuity between Gölcük and Akyazi coincides with the ruptures described from field survey. No deformation is visible out of this main fault zone. N-S offsets are below the noise level. Tectonic setting, with main active faults, of the  $M_s = 7.4$ , 17 August 1999 Izmit earthquake is given in inset. Focal mechanism is from USGS catalog, and epicenter (star) is from Kandilli Observatory. Shaded boxes denote location of SPOT and ERS images. Black lines show approximate location of ruptures described from field investigations. (b) Stack of all profiles across fault zone from Izmit to Sapanca lake. Offsets show a clear jump at the fault trace, indicating a fault zone  $<200$  m wide. Mean slip is  $3.5 \pm 0.2$  m. The two profiles differ due to uncertainty on SPOT changing attitude. Shaded line shows theoretical displacements derived from elastic modeling (Table 1). (c) Slip distribution between Gölcük and Akyazi measured from 1-km-wide swaths across the fault. Uncertainties are estimated from the variability of offsets on each side of the fault. The estimated slip at Sapanca lake is ascribed a large uncertainty because it is based on linear extrapolation of offsets across the lake. Field measurements are shown for comparison.





**Figure 2.** (opposite) Measured and modeled interferograms. (a) Differential SAR interferogram computed from ERS-2 SAR images (orbit frames 22556/815 and 23057/815) acquired on 13 August 1999 and 17 September 1999, respectively. A fringe represents 2.8 cm of ground shift away from satellite along the radar line of sight. Areas with a high level of noise result from temporal decorrelation. Near-fault areas are also poorly coherent due to either ground shaking or signal saturation. Interferogram is also affected by low-frequency noise (for wavelength larger than  $\sim 10$  km) with an amplitude of a few fringes induced by atmospheric artifacts and residual errors in determination of orbits [e.g., *Reilinger et al.*, 2000]. Orbital parameters tuned to remove a residual plane that cannot be explained by coseismic ground deformation. This linear correction amounts to three fringes in azimuth and four fringes in range. (b) Unwrapped interferogram. Phase unwrapping removes the  $2\pi$  ambiguity on measurements and yields relative displacements. Because the fault zone cuts across the interferogram, areas north and south of the fault zone are not related to the same reference frame, they were adjusted in order to fit elastic modeling. (c) Synthetic interferogram computed from the segmented fault model given in Table 1. We use an elastic half-space dislocation model and consider vertical faults that reach the surface. Geometry and mean slip along the Karamürsel-Gölcük, Gölcük-Izmit, Izmit-Sapanca, and Sapanca-Akyazi segments were constrained from the offsets. The fault extents at depth and dip angle were adjusted for each segment.