

A tsunami simulation of Hakata Bay using the viscous shallow-water equations

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Abstract The tsunami generated by the Great East Japan Earthquake caused serious damage to the coastal areas of the Tohoku district. Numerical simulations are used to predict damage caused by tsunamis. Shallow-water equations are generally used in numerical simulations of tsunami propagation from the open sea to the coast. This research focuses on viscous shallow-water equations and attempts to generate a computational method using finite element techniques based on the previous investigations of Kanayama and Ohtsuka (Coast Eng Jpn 21:157–171, 1978). First, the viscous shallow-water equation system is derived from the Navier–Stokes equations, based on the assumption of hydrostatic pressure in the direction of gravity. Next a numerical scheme is shown. Finally, tsunami simulations of Hakata Bay and Tohoku-Oki are shown using the approach.

Keywords Tsunami · Viscous shallow-water equations · Navier–Stokes equations · Finite element method

Mathematics Subject Classification 65

1 Introduction

The coastal areas of the Tohoku district suffered serious damage from the tsunami caused by the 2011 off the Pacific Coast of Tohoku Earthquake that occurred on March 11, 2011 [1]. Now, numerical simulations are used to develop disaster-prevention measures to deal with such tsunami disasters. They are used to predict potential future tsunami disasters, to design disaster-prevention facilities such as coastal breakwaters

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and levees, and in tsunami forecasts to predict tsunami attacks immediately after an earthquake occurs [2,3]. The Central Disaster Prevention Council [4] prepares a basic disaster prevention plan and participates in the determination of important disaster-prevention matters. If the Tonankai–Nankai Earthquake occurs, it may cause considerable damage. Consequently, the council has been performing numerical calculations to predict the wave height and the arrival time when the tsunami reaches the coast.

The waveform of a tsunami changes significantly in coastal areas compared with the open sea. This is because the water depth becomes shallow and the wave height becomes high when the tsunami reaches coastal areas. In addition, the wave height is amplified by the complex shapes of the closed-off sections of bays, as well as by the ocean floor topography. The shallow-water long-wave equations are generally used in the numerical simulation of tsunami propagation from the open sea to the coast [5,6]. In this study, we have focused on the viscous shallow-water equations [7] and attempted the computation by the finite element method based on the method proposed by Kanayama and Ohtsuka [8]. Thus far, we have studied the two-layer viscous shallow-water equations [9,10], obtained using the layer model [11], which is a well-known technique for quasi-three-dimensional computations. In these, we showed that the mass conservation law and the energy conservation law are satisfied under physically appropriate conditions [12]. We also found that the two-layer equations correspond to the one-layer equations [7] when the upper and lower layers have the same densities and the upper and lower layer equations are combined. Recently, we have found an interesting survey [13] which deals with sound mathematical topics related to the viscous shallow-water equations.

In this paper, to let it be self-contained, the viscous shallow-water equations are again derived from the Navier–Stokes equations in which the hydrostatic pressure in the direction of gravity is assumed. In the numerical analysis of tsunamis, the viscosity term is often omitted or simply added [5]. In this study, however, a computational method that does not omit the viscosity term is adopted. That is because it enables more rigorous analysis to be performed and we intend to include the viscosity term in future stress analysis of tsunamis. The approximation scheme is given below and simulation results for a Hakata Bay model and a Tohoku-Oki model are presented as computational examples.

In the computational example of Hakata Bay, the tsunami generating area is taken to be the epicenter of the West off Fukuoka Prefecture Earthquake, which occurred in 2005 in the Sea of Genkai located northwest offshore of Fukuoka Prefecture. The validity of this computational method can be confirmed by examining the tsunami propagation velocity based on the numerical analysis results. It can also be confirmed that when the tsunami reaches land in the coastal areas, waves are reflected with islands functioning as breakwaters, suppressing the wave height in our computation.

2 Derivation of the viscous shallow-water equations

As shown in Fig. 1, we consider the shallow-water long-wave flow in which the wavelength is sufficiently long relative to the water depth [7]. First, the viscous shallow-water equations are derived. Orthogonal coordinates [m] are used, where x_1 and x_2

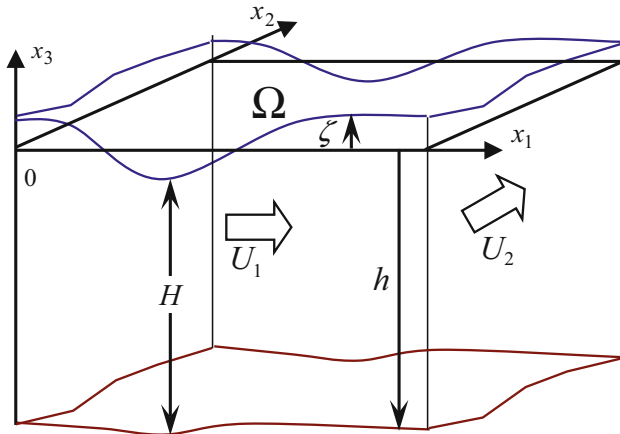


Fig. 1 Variables of the computational model

represent directions in the horizontal plane and x_3 represents the vertical direction, and the time [s] is represented by t . Assuming the hydrostatic pressure in the x_3 direction, the following Navier–Stokes equations are used with the external forces, namely the Coriolis forces and the gravity are assumed to act in the x_1 and x_2 directions and the x_3 direction, respectively.

$$\sum_{j=1}^3 \frac{\partial u_j}{\partial x_j} = 0, \tag{1}$$

$$\frac{\partial u_1}{\partial t} + \sum_{j=1}^3 u_j \frac{\partial u_1}{\partial x_j} = -\frac{1}{\rho} \frac{\partial p}{\partial x_1} + \frac{1}{\rho} \sum_{j=1}^3 \frac{\partial \tau_{1j}}{\partial x_j} + f u_2, \tag{2}$$

$$\frac{\partial u_2}{\partial t} + \sum_{j=1}^3 u_j \frac{\partial u_2}{\partial x_j} = -\frac{1}{\rho} \frac{\partial p}{\partial x_2} + \frac{1}{\rho} \sum_{j=1}^3 \frac{\partial \tau_{2j}}{\partial x_j} + f u_1, \tag{3}$$

$$0 = -\frac{1}{\rho} \frac{\partial p}{\partial x_3} - g. \tag{4}$$

$u(x_1, x_2, x_3, t)$ represents the fluid velocity [m/s] in the x_i ($i = 1-3$) direction, $p(x_1, x_2, x_3, t)$ denotes the pressure [N/m²], ρ is the density [kg/m³], τ_{ij} is the stress [N/m²] in the x_i direction acting on the x_j plane, f is the Coriolis coefficient [1/s], and g is the acceleration [m/s²] due to the gravity. In addition, the ordinate [m] of the water surface (water level) is represented by ζ and the ordinate [m] of the bottom surface is represented by h .

Equations (1)–(3) are integrated with respect to x_3 from the bottom h to the water surface ζ . The velocity component in the normal direction is assumed to be zero at the water and bottom surfaces. By doing this, the viscous shallow-water equations expressed by the averaged velocity can be derived. If the layer thickness is $H(x_1, x_2, t)$ and the average velocity in the x_i ($i = 1-2$) direction is $U_i(x_1, x_2, t)$, then H and U_i are given by the following equations:

$$H = \zeta - h, \quad U_i = \frac{1}{H} \int_h^\zeta u_i dx_3. \tag{5}$$

The fixed density of the layer is represented by ρ and the stress in the x_i direction acting on the x_j plane is represented by $\tilde{\tau}_{ij}$. The horizontal viscosity constant [N s/m²] is represented by μ_H , the wind effect coefficient is represented by θ , the air density is represented by ρ_a , the wind speed in the x_i direction is represented by W_i , and the Chezy coefficient [m^{1/2}/s] is represented by C . Then, the following viscous shallow-water equations are derived:

$$\frac{\partial \zeta}{\partial t} + \sum_{j=1}^2 \frac{\partial}{\partial x_j} (H U_j) = 0, \tag{6}$$

$$H \left(\frac{\partial U_i}{\partial t} + \sum_{j=1}^2 U_j \frac{\partial U_i}{\partial x_j} \right) = -gH \frac{\partial \zeta}{\partial x_i} + \frac{1}{\rho} \sum_{j=1}^2 \frac{\partial}{\partial x_j} (H \tilde{\tau}_{ij}) + \frac{1}{\rho} \theta \rho_a W_i ((W_1)^2 + (W_2)^2)^{\frac{1}{2}} - \frac{g}{C^2} U_i ((U_1)^2 + (U_2)^2)^{\frac{1}{2}} + (-1)^{i+1} f H U_{i+1}, \tag{7}$$

where

$$\tilde{\tau}_{ij} = \mu_H \left(\frac{\partial U_i}{\partial x_j} + \frac{\partial U_j}{\partial x_i} \right), \quad U_3 = U_1. \tag{8}$$

Finite element approximation is performed for the viscous shallow-water equations (6) and (7) with given initial conditions and boundary conditions.

The computational domain is the two-dimensional polygonal region Ω surrounded by the boundaries Γ_c and Γ_o . In this region, the orthogonal coordinates $x = (x_1, x_2)$ are used. The boundary conditions on Γ_c and Γ_o and the initial conditions at $t = 0$ are as follows:

Boundary conditions:

$$U_i(x, t) = 0 \quad \text{on } \Gamma_c, \tag{9}$$

$$\left. \begin{aligned} \zeta(x, t) &= \zeta_{\Gamma_o}(x, t) \\ \sum_{j=1}^2 \tilde{\tau}_{ij} n_j &= 0 \end{aligned} \right\} \quad \text{on } \Gamma_o. \tag{10}$$

Here, n_i is the component in the x_i direction of the unit normal vector on the boundary. On Γ_o , $\zeta_{\Gamma_o}(x, t)$ is specified.

Initial conditions:

$$U_i(x, 0) = U_{i0}(x), \quad \zeta(x, 0) = \zeta_0(x). \tag{11}$$

Here, $U_{i0}(x)$ and $\zeta_0(x)$ are the initial values of U_i and ζ , respectively.

3 Finite element approximations for the viscous shallow-water equations

The finite element approximation [8] is as follows. First, Eqs. (6) and (7) are multiplied by test functions and then integrated over the computational domain Ω . Subsequently, the approximation is carried out for terms containing a derivative with respect to time by using an explicit method. For terms containing spatial derivatives, the finite element approximation is carried out by using the piecewise linear basis function $\hat{\varphi}_k$. For terms without spatial derivatives, the approximation is carried out by using the corresponding step function $\bar{\varphi}_k$. The step function $\bar{\varphi}_k$ is 1 in the barycentric region around the node k and is 0 at other places.

$$\begin{aligned}
 & \left(\frac{\bar{\zeta}^{n+1} - \bar{\zeta}^n}{\Delta t}, \bar{\varphi}_k \right) + \left(\sum_{j=1}^2 \frac{\partial(\hat{H}^n \hat{U}_j^n)}{\partial x_j}, \hat{\varphi}_k \right) = 0. \tag{12} \\
 & \left(\bar{H}^{n+1} \left(\frac{\bar{U}_i^{n+1} - \bar{U}_i^n}{\Delta t} \right), \bar{\varphi}_k \right) + \left(\hat{H}^{n+1} \sum_{j=1}^2 \hat{U}_j^n \frac{\partial \hat{U}_i^n}{\partial x_j}, \hat{\varphi}_k \right) \\
 & = - \left(g \hat{H}^{n+1} \frac{\partial \hat{\zeta}^{n+1}}{\partial x_i}, \hat{\varphi}_k \right) - \sum_{j=1}^2 \left(\frac{1}{\rho} \hat{H}^{n+1} \mu_H \left(\frac{\partial \hat{U}_i^n}{\partial x_j} + \frac{\partial \hat{U}_j^n}{\partial x_i} \right), \frac{\partial \hat{\varphi}_k}{\partial x_j} \right) \\
 & + \left(\frac{1}{\rho} \theta \rho_a W_i ((W_1)^2 + (W_2)^2)^{\frac{1}{2}}, \bar{\varphi}_k \right) - \left(\frac{g}{C^2} \bar{U}_i^n ((\bar{U}_1^n)^2 + (\bar{U}_2^n)^2)^{\frac{1}{2}}, \bar{\varphi}_k \right) \\
 & + ((-1)^{i+1} f \bar{H}^{n+1} \bar{U}_{i+1}^n, \bar{\varphi}_k). \tag{13}
 \end{aligned}$$

In the above, ζ_k^{n+1} is determined from (12). From the obtained value and $U_{i,k}^n$, the value $U_{i,k}^{n+1}$ is determined by (13). Here, ζ_k^n and $U_{i,k}^n$ are approximate values of $\zeta(x, t)$ and $U(x, t)$ respectively, at the node k after n time steps, and Δt represents the size of time steps. In addition, the following symbols are also used:

$$\left(\frac{\partial U_i}{\partial x_j}, v \right) = \int_{\Omega} \frac{\partial U_i}{\partial x_j} v \, dx, \tag{14}$$

$$\hat{U}_i^n = \sum_k U_{i,k}^n \hat{\varphi}_k, \bar{U}_i^n = \sum_k U_{i,k}^n \bar{\varphi}_k \tag{15}$$

Because of the presence of the basic boundary conditions, (12) holds for all nodes except for the nodes on Γ_o and (13) holds for all the nodes except for the nodes on Γ_c . However, on the boundary Γ_o , an approximation method is adopted in which the advective term of (13) uses Tabata’s upwind approximation [8, 14]. A mathematical justification for the approximation scheme for linearized equations related to the above scheme was given by Kanayama and Ushijima [15, 16] (see also Appendix A for self-containedness).

4 Computational examples

The numerical results for Hakata Bay are shown. The tsunami generating area was taken to be the epicenter of the West off Fukuoka Prefecture Earthquake, which occurred in the Sea of Genkai located northwest offshore of Fukuoka Prefecture on March 20, 2005 (maximum seismic intensity: 6 lower). Because it was a strike slip-fault earthquake, no tsunami was observed despite the hypocenter being on the ocean floor. It is therefore noted that our numerical results can not be compared with observed data.

The computation is performed with a Coriolis coefficient $f = 2\omega \sin \phi$, a planetary angular velocity $\omega = 0.00416$ [deg/s], a latitude $\phi = 35$ [deg], a horizontal viscosity coefficient $\mu_H = 0.001$ [N s/m²], a wind effect coefficient $\theta = 0.0026$, a wind speed $W_1 = W_2 = 0$ [m/s], a Chezy coefficient $C = 61.4$ [m^{1/2}/s], and a time step size $\Delta t = 0.1$ [s]. This time step size is found to be too small, because we can get similar results even if a 10 times larger time step size is used for the following mesh data. Details will later be published, referring to the linear stability analysis in Appendix A. Though the full nonlinear stability analysis is very difficult, we believe that the present linear stability condition is useful. In fact, the similar stability condition of [18] in the case of $\theta = 1$ produces a good estimate of the 10 times larger time step size (1 [s]) in the numerical experiments. Regarding the mesh data, the total number of nodes is 11,370, the total number of elements is 22,099, and the ocean floor ordinate is fixed at $h = -10$ [m]. This constant depth is also not a strong simplification, because the actual depth in Hakata Bay is almost flat. We set the initial conditions $\zeta_0(x) = 0$ and $U_{i0}(x) = 0$, and the boundary conditions $\zeta_{\Gamma_0}(x, t) = 5$ [m] ($0 < t < 60$ [s]) at the tsunami generating area and $\zeta_{\Gamma_0}(x, t) = 0$ [m] at other areas. See Fig. 2 where the tsunami generating line is shown in different color (red). In general, the tsunami is excited by two ways. The first one is to consider it in the initial condition of the water surface ζ , for which we do not have sufficient input information in this artificial tsunami of Hakata Bay. The second one is to consider it in the boundary condition of ζ as in this paper. In our setting, the computational domain is not so wide that the above approach may be the only way.

Figure 3a–f show the ordinates of the water surface (water level) every 4 [min] from after 4 [min] to after 24 [min]. It is apparent that the tsunami propagates to the closed-off section of the bay. In addition, since inundation of the tsunami is not taken into consideration in our simulation, part of the tsunami is reflected by the island.

Figure 4 shows the water level change at the three points A–C in Fig. 5. It is confirmed that the presence of an island near the coast suppresses the wave height on the land side of the island. The distance from the tsunami generation area to the point A is 15.3 [km], and the arrival time of the first wave at the point A is 19 [min]. If the propagation velocity is calculated from the arrival time of the first wave of a tsunami, the obtained value (48 [km/h] = 15.3 [km]/ 19 [min]) is somewhat larger than the simplified theoretical value $\sqrt{g|h|}$ (≈ 36 [km/h]) where $|h|$ is the water depth. This is considered to be due to the surrounding land.

Next, the numerical results for Tohoku-Oki are shown. When not specifically defined, the same physical values as for Hakata Bay were used. Regarding the mesh

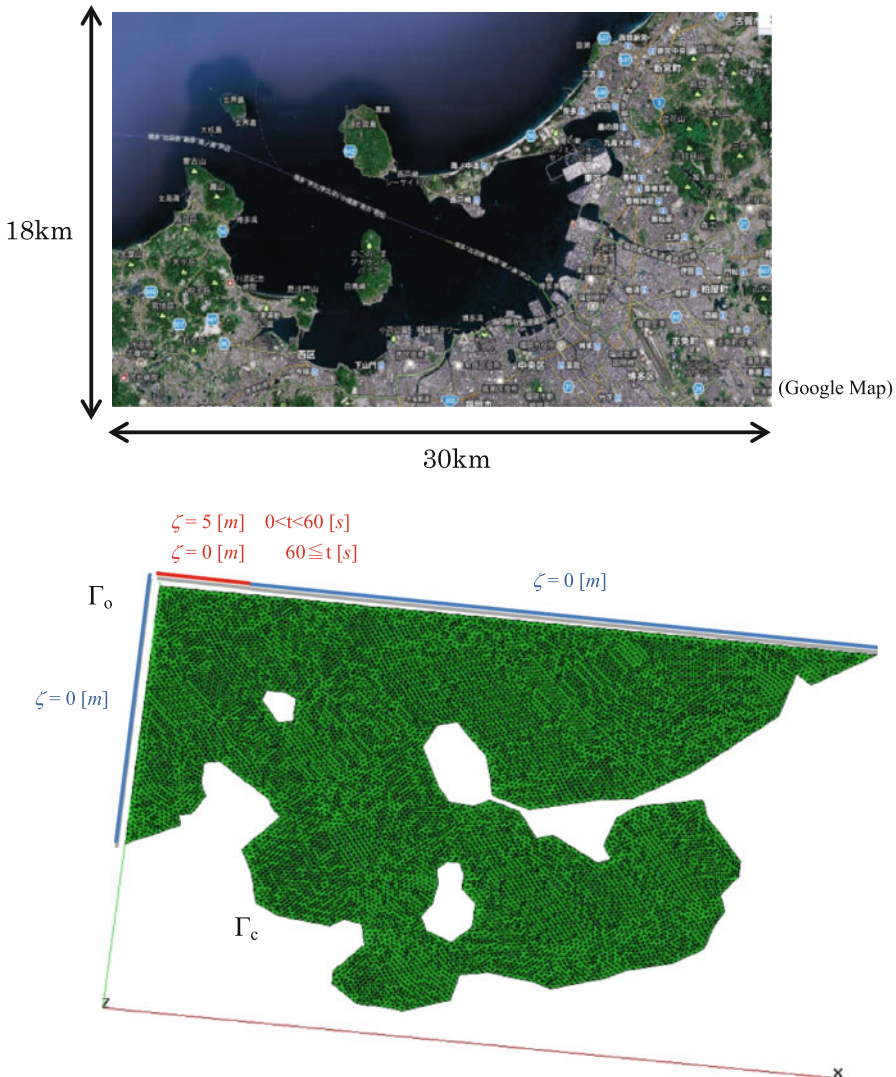


Fig. 2 A computational mesh of Hakata Bay (color figure online)

data, the total number of nodes is 56,562, the total number of elements is 113,013, and the ocean floor ordinate is set to become deep gradually from $h = -10 [m]$ to $h = -1,000 [m]$. We set the initial conditions at $\zeta_0(x) = 0$ and $U_{i0}(x) = 0$, and the boundary conditions $\zeta_{\Gamma_0}(x, t) = 50 [m]$ ($0 < t < 60 [s]$) at the tsunami generating area and $\zeta_{\Gamma_0}(x, t) = 0 [m]$ at other areas. See Fig. 6. It is noted that 50 [m] at Γ_0 may be too high. In our computation, the tsunami arrived at Oshika Peninsula in Fig. 6 after 20 min and the highest wave height reached 15 [m]. These numerical results should more carefully be checked with boundary data on Γ_0 . Numerical results with suitable initial data for the tsunami generation will later be published.

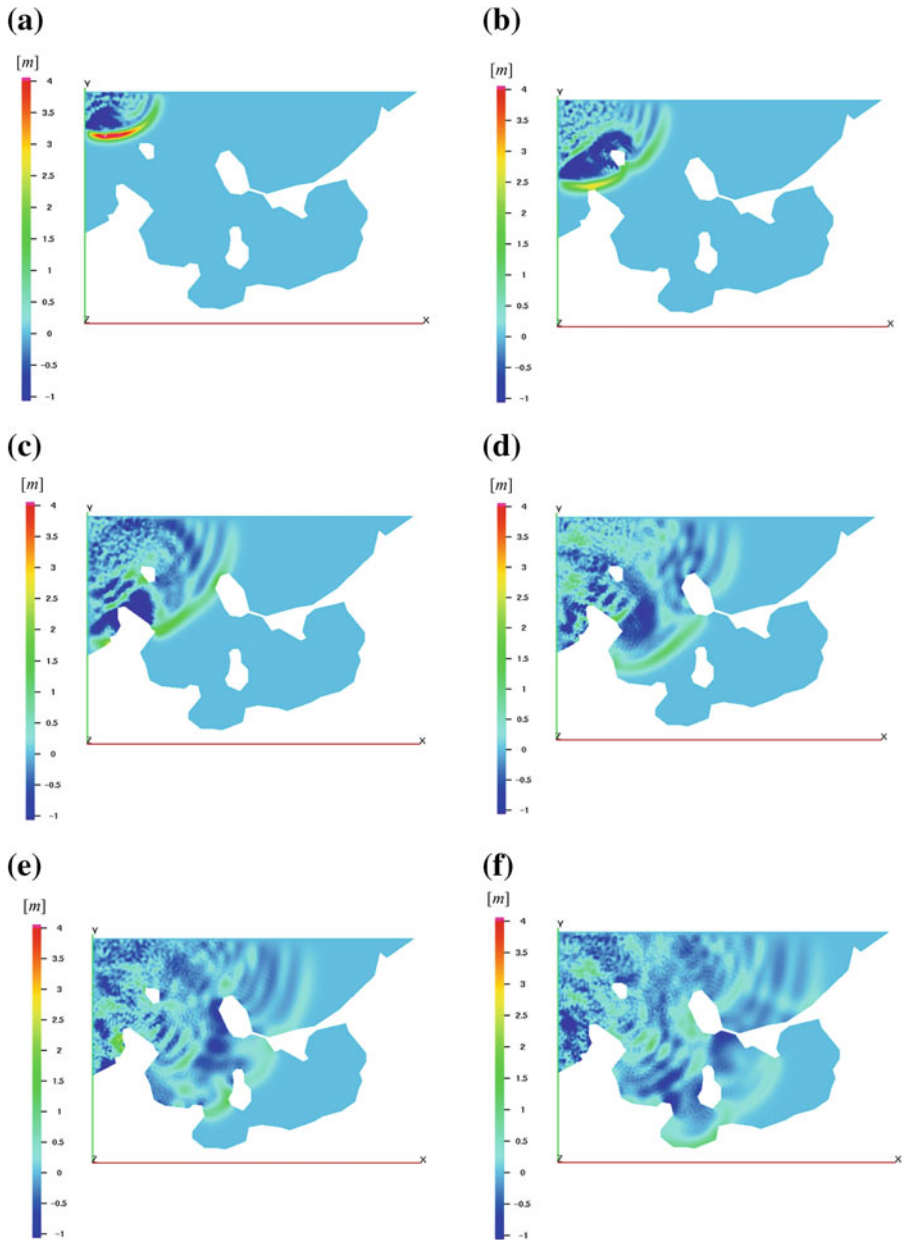


Fig. 3 **a** Contour map of ζ after 4 min. **b** Contour map of ζ after 8 min. **c** Contour map of ζ after 12 min. **d** Contour map of ζ after 16 min. **e** Contour map of ζ after 20 min. **f** Contour map of ζ after 24 min

Figure 7 shows the ordinates of the water surface (water level) after 18 [min]. Since the computational domain is not wide, it looks that there is a reflection from the northern boundary in Fig. 7. This artificial reflection can be removed by suitable boundary conditions on Γ_0 . Details will later be published. Figure 8 shows the water

Fig. 4 Time histories of ζ

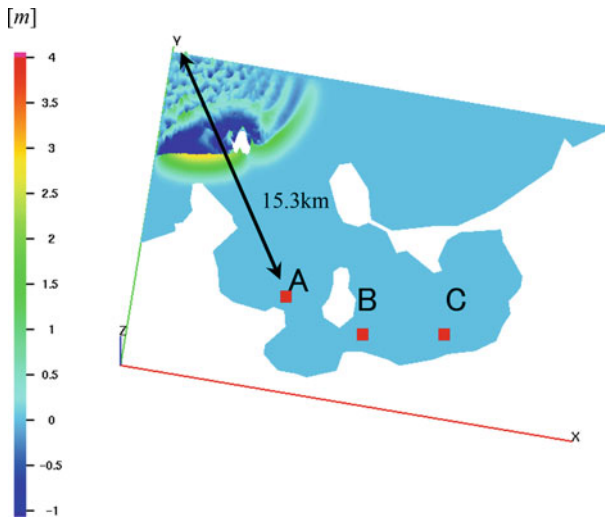
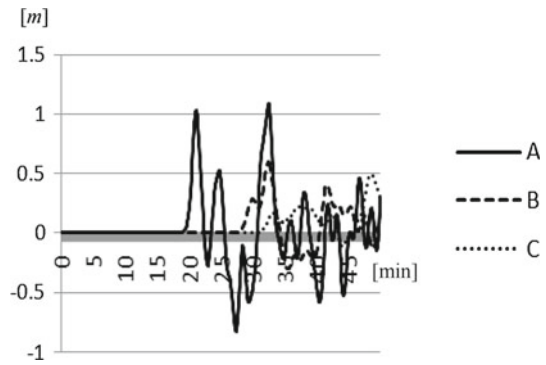


Fig. 5 Contour map of ζ

level change at the two points A and B in Fig. 6. The wave height at the point B is higher than the point A after about 1,500 [s]. When the tsunami reaches coastal areas, the water depth becomes shallow and the wave height becomes high.

5 Concluding remarks

In this study, the viscous shallow-water equations have again been derived from the Navier–Stokes equations in which the hydrostatic pressure in the direction of gravity is assumed. Tsunami propagation is then simulated by a finite element computation. Using the present Hakata Bay model, in which the tsunami generating area is taken to be the epicenter of the West off Fukuoka Prefecture Earthquake (2005), tsunami propagation from the open sea to the coastal area may be produced. In addition, we have constructed a system in which the tsunami arrival time and the wave height at the time of tsunami attack can be obtained from numerical results. Since our analysis takes the viscosity term into account, this study can be considered to be a preliminary

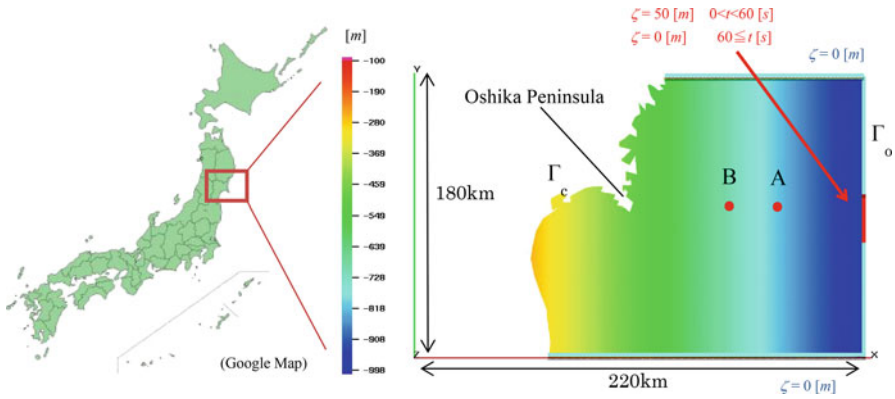


Fig. 6 A computational model of Tohoku-Oki

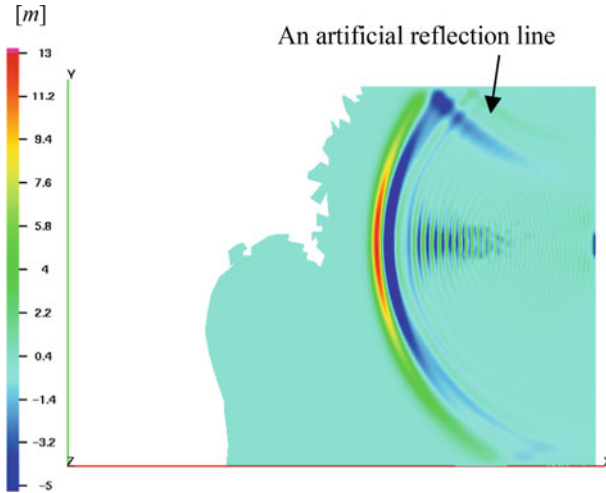


Fig. 7 Contour map of ζ

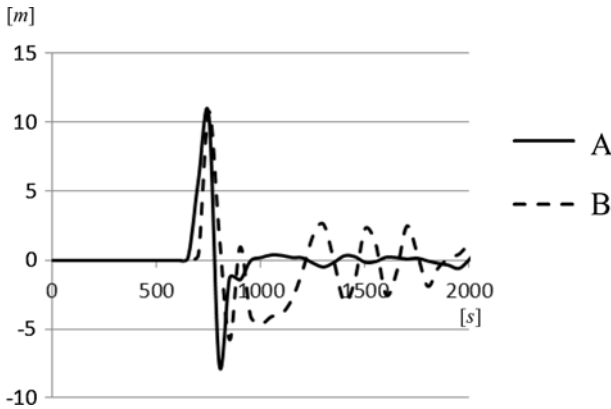
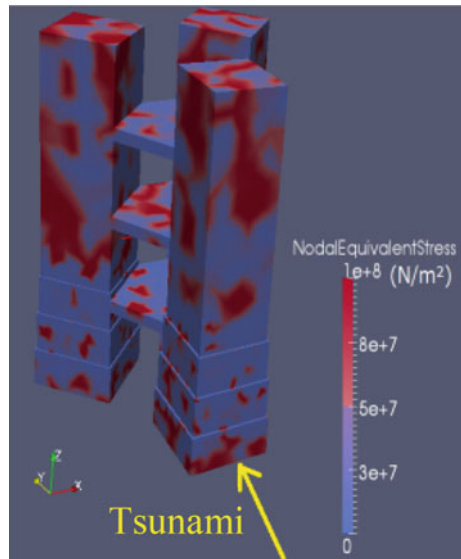


Fig. 8 Time histories of ζ

Fig. 9 Structural analysis in the case of tsunami attack



study for future planned stress analyses of tsunamis. See Fig. 9. We want to check whether the hydrostatic pressure is always the dominant force in the case of tsunami attack, compared with other forces including the viscous force.

Appendix A

A.1 Introduction

This appendix summarizes contents of Kanayama and Ushijima [15, 16] for self-containedness. First we introduce a linearized system [15] of the viscous shallow-water equations. After the derivation of the model problem, we reformulate the linear initial-boundary value problem within the framework of Hilbert space theory. Then, the existence and uniqueness of the solution of the reformulated problem follows from Hille–Yosida theorem. Next, this appendix concerns the finite element analysis of linearized viscous shallow-water equations [16]. Our scheme is obtained by the use of the “piecewise-linear” continuous finite element approximation for both the water-level and the velocity with respect to the space variables, together with the step by step integration method for the time variable. The stability criterion of the scheme and the convergence of the approximate solutions to the exact one are obtained in L^2 sense.

A.2 The derivation of a linearized viscous shallow-water system

Let us consider the following nonlinear stationary problem

$$\begin{aligned}
 &U_0 \cdot \nabla U_0 + [f]U_0 + g\nabla\zeta_0 \\
 &= \frac{1}{H_0\rho} \left\{ \operatorname{div}(H_0\tau_0) + \theta\rho_a|V|V - \frac{\rho g}{C^2}|U_0|U_0 \right\} \quad \text{in } \Omega, \tag{16} \\
 &\operatorname{div}(H_0U_0) = 0 \quad \text{in } \Omega. \tag{17}
 \end{aligned}$$

Supposing that the domain Ω is an open bounded set in the plane R^2 with the C^∞ -class boundary Γ , we set the following boundary condition:

$$U_0 = 0 \quad \text{on } \Gamma. \tag{18}$$

We assume the existence of a sufficiently smooth solution pair $\{U_0, \zeta_0\}$ of the above problem (16)–(18). Further, let us suppose the existence of two positive constants \overline{H}_0 and \underline{H}_0 such that

$$0 < \underline{H}_0 \leq H_0 = \zeta_0 - h \leq \overline{H}_0 \quad \text{in } \Omega. \tag{19}$$

Furthermore we assume $U_0 \in (C^\infty(\overline{\Omega}))^2$ and $H_0 \in C^\infty(\overline{\Omega})$, where $\overline{\Omega}$ denotes the closure of Ω , and $C^\infty(\overline{\Omega})$ consists of all infinitely differentiable functions on $\overline{\Omega}$. In the following, $C(\overline{\Omega})$ denotes the totality of continuous functions on $\overline{\Omega}$.

As a model linear initial-boundary value problem, we adopt the following linearized viscous shallow-water equations (20) and (21):

$$\frac{\partial U}{\partial t} + U_0 \cdot \nabla U + [f]U + g\nabla\zeta = \frac{1}{H_0\rho} \left\{ \operatorname{div}(H_0\tau) - \frac{\rho g}{C^2}|U_0|U \right\} \quad \text{in } \Omega, \tag{20}$$

$$\frac{\partial \zeta}{\partial t} + \operatorname{div}(H_0U) = 0 \quad \text{in } \Omega, \tag{21}$$

with the initial and boundary conditions:

$$U(x, 0) = \tilde{U}(x) \quad \text{in } \Omega, \tag{22}$$

$$\zeta(x, 0) = \tilde{\zeta}(x) \quad \text{in } \Omega, \tag{23}$$

$$U(x, t) = 0 \quad \text{on } \Gamma, \tag{24}$$

where $\tilde{U}(x)$ and $\tilde{\zeta}(x)$ are sufficiently smooth initial data. In the sequel, conditions (17), (18) and (19) will be assumed, and (16) will not be used.

A.3 The unique solvability of the linear initial-boundary value problem

Let $L^2(\Omega)$ be the space of functions which are square integrable on Ω . The space $(L^2(\Omega))^3$ is considered as the product space of $(L^2(\Omega))^2$ and $L^2(\Omega)$, and any $v \in (L^2(\Omega))^3$ is represented as $v = \begin{pmatrix} V \\ \eta \end{pmatrix}$ with $V \in (L^2(\Omega))^2$ and $\eta \in L^2(\Omega)$.

Introduce the following inner product $(\mathbf{u}, \mathbf{v})_X$ in $(L^2(\Omega))^3$:

$$(\mathbf{u}, \mathbf{v})_X = (H_0U, V)_{(L^2(\Omega))^2} + g(\zeta, \eta)_{L^2(\Omega)} \quad \text{for } \mathbf{u} = \begin{pmatrix} U \\ \zeta \end{pmatrix}, \mathbf{v} = \begin{pmatrix} V \\ \eta \end{pmatrix} \in (L^2(\Omega))^3.$$

Due to (19), $(L^2(\Omega))^3$ becomes a real Hilbert space with this inner product $(\cdot, \cdot)_X$, which is denoted by X in the sequel.

Let us consider the following X -valued evolution problem:

$$\frac{d\mathbf{u}}{dt} = \mathbf{A}\mathbf{u} \quad (t > 0), \tag{25}$$

$$\mathbf{u}(0) = \mathbf{a} \in D(\mathbf{A}), \tag{26}$$

where the linear operator \mathbf{A} is defined by

$$\mathbf{A}\mathbf{u} = \begin{bmatrix} -U_0 \cdot \nabla U - [f]U + \frac{1}{H_0\rho} \left\{ \operatorname{div}(H_0\tau) - \frac{\rho g}{C^2} |U_0|U \right\} - g\nabla\zeta \\ -\operatorname{div}(H_0U) \end{bmatrix}, \tag{27}$$

with the domain $D(\mathbf{A})$ given by

$$D(\mathbf{A}) = \begin{matrix} (H_0^1(\Omega) \cap H^2(\Omega))^2 \\ \times \\ H^1(\Omega). \end{matrix} \tag{28}$$

Here, the standard notations of Sobolev spaces are used. Namely, let m be a non-negative integer. The Sobolev space $H^m(\Omega)$ of order m on Ω is defined by $H^m(\Omega) = \{v : \partial^\alpha v \in L^2(\Omega), |\alpha| \leq m\}$, where

$$\partial^\alpha v = \frac{\partial^{\alpha_1 + \alpha_2} v}{\partial x_1^{\alpha_1} \partial x_2^{\alpha_2}}, \quad \alpha = \{\alpha_1, \alpha_2\}, \quad |\alpha| = \alpha_1 + \alpha_2.$$

We consider $H^m(\Omega)$ the Hilbert space with the norm:

$$\|v\|_m = \left(\sum_{|\alpha| \leq m} \int_{\Omega} |\partial^\alpha v|^2 dx \right)^{1/2}.$$

Let $\mathcal{D}(\Omega)$ be the totality of infinitely differentiable functions on Ω with compact supports in Ω . The closure of $\mathcal{D}(\Omega)$ in $H^m(\Omega)$ is denoted by $H_0^m(\Omega)$.

Theorem 1 *The linear operator \mathbf{A} is a densely defined dissipative operator, that is,*

- (i) $D(\mathbf{A})$ is dense in X ,
- (ii) $(\mathbf{A}\mathbf{u}, \mathbf{u})_X \leq 0$ for any $\mathbf{u} \in D(\mathbf{A})$.

The range $R(A-\lambda)$ is dense in X . Hence, the closure \bar{A} of A exists and is a maximal dissipative operator in X , since A is dissipative by Theorem 1. \bar{A} generates a contraction semi-group of linear operators $\{e^{t\bar{A}}\}_{t \geq 0}$ acting in X . Therefore, we arrive at the following theorem [15].

Theorem 2 Consider the X -valued evolution problem :

$$\begin{cases} \frac{du}{dt} = \bar{A}u & (t > 0), \\ u(0) = a \in D(\bar{A}). \end{cases} \tag{E}$$

Then, (E) has a unique solution $u(t) = e^{t\bar{A}}a$ satisfying

$$\|u(t)\|_X \leq \|a\|_X \quad (t \geq 0).$$

Remark 1 For any $a \in X$, the function $u(t) = e^{t\bar{A}}a$ is called the generalized solution of the problem (E).

A.4 A finite element semi-discrete approximation

For simplicity of notations, we introduce bilinear forms $a(U, V), b(U, V), c(\zeta, V)$ and $d(U, \eta)$ for $U, V \in (H_0^1(\Omega))^2$ and $\zeta, \eta \in H^1(\Omega)$ as follows:

$$\begin{cases} a(U, V) = \frac{1}{\rho}(H_0\tau, \nabla V)_{(L^2(\Omega))^4} + \frac{g}{C^2}(|U_0|U, V)_{(L^2(\Omega))^2}, \\ b(U, V) = (U_0 \cdot \nabla U, H_0V)_{(L^2(\Omega))^2} + ([f]U, H_0V)_{(L^2(\Omega))^2}, \\ c(\zeta, V) = g(\nabla\zeta, H_0V)_{(L^2(\Omega))^2}, \\ d(U, \eta) = g(\operatorname{div}(H_0U), \eta)_{(L^2(\Omega))}. \end{cases} \tag{29}$$

Define the function space V and the bilinear form $\mathbf{a}(u, v)$ for $u, v \in V$ as follows:

$$\begin{cases} V = (H_0^1(\Omega))^2 \times H^1(\Omega), \\ \mathbf{a}(u, v) = -\{a(U, V) + b(U, V) + c(\zeta, V) + d(U, \eta)\} \\ \text{for } u = \begin{pmatrix} U \\ \zeta \end{pmatrix}, v = \begin{pmatrix} V \\ \eta \end{pmatrix} \in V. \end{cases} \tag{30}$$

Set the following weak formulation (π) of (E):

$$\begin{cases} \text{Find } u(t) \in V (t \geq 0) \text{ such that} \\ \frac{d}{dt}(u(t), v)_X = \mathbf{a}(u(t), v) \text{ for any } v \in V \text{ and } t > 0, \\ u(0) = a \in V. \end{cases} \tag{\pi}$$

Let \mathfrak{T}_h be a ‘‘triangulation’’ of the domain Ω with the representative length h . The curved elements due to Zlamal [17] are adopted near the boundary Γ , while the triangular elements are used in interior of Ω . We assume that the family of ‘‘triangulation’’

$\{\mathfrak{S}_h : 0 < h \leq \bar{h} < \infty\}$ is regular, and satisfies the inverse assumption. Namely, there are positive constants Σ and Λ independent of h such that

$$\frac{h_T}{\rho_T} \leq \Sigma, \quad \frac{h}{h_T} \leq \Lambda \quad \text{for } T \in \mathfrak{S}_h,$$

where h_T and ρ_T denote the diameters of an element T and the inscribed circle of T , respectively.

Let W_h be the “piecewise-linear” continuous finite element space constructed by \mathfrak{S}_h :

$$W_h = \{\eta_h \in C(\bar{\Omega}) : \eta_h \text{ is a “polynomial” with degree at most 1 on each } T \in \mathfrak{S}_h\}.$$

More precisely, $\eta_h \in W_h$ is a polynomial with degree at most 1 on each triangular element of \mathfrak{S}_h , while on a curved element T , $\eta_h(F_T(\hat{x}))$ is a polynomial with degree at most 1 on a reference triangle \hat{T} , where $F_T : \hat{T} \ni \hat{x} \rightarrow x = F_T(\hat{x}) \in T$ is the isomorphism from \hat{T} onto T determined by Zlamal’s method [17]. Now, define the vector-valued finite element space \mathcal{V}_h by

$$\mathcal{V}_h = \{V_h \in (W_h)^2 : V_h = 0 \text{ on } \Gamma\},$$

and let

$$V_h = \mathcal{V}_h \times W_h.$$

Then

$$V_h \subset V = (H_0^1(\Omega))^2 \times H^1(\Omega).$$

The following problem (π_h) is said to be the semi-discrete Galerkin approximation of (π) :

$$\begin{cases} \text{Find } \mathbf{u}_h(t) \in V_h \text{ such that} \\ \frac{d}{dt}(\mathbf{u}_h(t), \mathbf{v}_h)_X = \mathbf{a}(\mathbf{u}_h(t), \mathbf{v}_h) \text{ for any } \mathbf{v}_h \in V_h \text{ and } t > 0, \\ \mathbf{u}_h(0) = \mathbf{a}_h \in V_h. \end{cases} \quad (\pi_h)$$

The set V_h , being considered as a closed subspace of X , is denoted by X_h , which is a Hilbert space with the inner product $(\cdot, \cdot)_X$. Let A_h be the bounded linear operator acting on X_h , defined by the relation:

$$(A_h \mathbf{u}_h, \mathbf{v}_h)_X = \mathbf{a}(\mathbf{u}_h, \mathbf{v}_h) \text{ for any } \mathbf{u}_h, \mathbf{v}_h \in X_h.$$

Since we have (3.5)–(3.9) in [15] for $\mathbf{u} = \begin{pmatrix} U \\ \zeta \end{pmatrix} \in V$, it holds that

$$\mathbf{a}(\mathbf{v}, \mathbf{v}) \leq 0 \quad \text{for } \mathbf{v} \in V. \tag{31}$$

This estimate together with the finite dimensionality of X_h implies that the operator A_h is maximal dissipative. Hence, the exponential function $\{e^{tA_h} : t \geq 0\}$ forms a

contraction semi-group acting on X_h . The problem (π_h) is equivalently represented in the following X_h -valued evolution equation (E_h) :

$$\begin{cases} \frac{d\mathbf{u}_h}{dt} = \mathbf{A}_h \mathbf{u}_h & (t > 0), \\ \mathbf{u}_h(0) = \mathbf{a}_h \in X_h. \end{cases} \tag{E_h}$$

The unique solution of (E_h) is given by $\mathbf{u}_h(t) = e^{t\mathbf{A}_h} \mathbf{a}_h$, which satisfies

$$\|\mathbf{u}_h(t)\|_X \leq \|\mathbf{a}_h\|_X \quad (t \geq 0). \tag{32}$$

The main conclusion of this section [15] is:

Theorem 3 *Let $\mathbf{u}(t)$ and $\mathbf{u}_h(t)$ be the generalized solution of (E) and the solution of (E_h) , respectively. If $\lim_{h \rightarrow 0} \|\mathbf{a}_h - \mathbf{a}\|_X = 0$, then $\lim_{h \rightarrow 0} \|\mathbf{u}_h(t) - \mathbf{u}(t)\|_X = 0$ uniformly in $t \in [0, T]$ for any finite T .*

A.5 A class of finite difference schemes in time

First, several notations of spaces and operators are prepared to state our schemes. Let X be the Hilbert space $(L^2(\Omega))^2$ with the inner product:

$$(U, V)_X = (H_0 U, V)_{(L^2(\Omega))^2}, \tag{33}$$

and let Z be the space $L^2(\Omega)$ with the inner product:

$$(\zeta, \eta)_Z = g(\zeta, \eta)_{L^2(\Omega)}. \tag{34}$$

The spaces \mathcal{V}_h and W_h are denoted by X_h and Z_h when they are considered as the subspaces of X and Z , respectively. The product Hilbert space $\begin{matrix} X_h \\ \times \\ Z_h \end{matrix}$ is denoted by X_h . The bounded operators $A_h \in L(X_h)$, $B_h \in L(X_h)$, $C_h \in L(Z_h, X_h)$ and $D_h \in L(X_h, Z_h)$ are defined so as to satisfy

$$\begin{cases} (A_h U_h, V_h)_{X_h} = a(U_h, V_h), \\ (B_h U_h, V_h)_{X_h} = b(U_h, V_h), \\ (C_h \zeta_h, V_h)_{X_h} = c(\zeta_h, V_h), \\ (D_h U_h, \eta_h)_{Z_h} = d(U_h, \eta_h), \\ \text{for any } U_h, V_h \in X_h \text{ and } \zeta_h, \eta_h \in Z_h. \end{cases} \tag{35}$$

Now our finite difference schemes are given in the following :

$$\left\{ \begin{array}{l} \text{Find } \mathbf{u}_n = \begin{pmatrix} U_n \\ \zeta_n \end{pmatrix} \in \mathbf{X}_h = \begin{matrix} X_h \\ \times \\ Z_h \end{matrix} \\ \text{for } n = 1, 2, \dots \text{ such that} \\ \frac{U_n - U_m}{\tau} + A_h U_m + B_h(\theta U_m + (1 - \theta)U_n) + C_h \zeta_m = 0, \\ \frac{\zeta_n - \zeta_m}{\tau} + D_h U_n = 0, \\ \text{with } n = m + 1, m = 0, 1, 2, \dots, \end{array} \right. \quad (\mathbf{E}_{h,\tau})$$

where τ is a given positive time increment, θ is a parameter satisfying

$$0 \leq \theta \leq 1, \tag{36}$$

and the initial value $\mathbf{u}_0 = \begin{pmatrix} U_0 \\ \zeta_0 \end{pmatrix}$ is given in \mathbf{X}_h . It is noted that Ushijima [18] deals with the case of $\theta = 1$, while Kanayama–Ushijima [19] slightly modifies the case of $\theta = \frac{1}{2}$.

Introduce $a_i(U, V)$ and $b_i(U, V)$ ($i = 0, 1$) as follows:

$$\left\{ \begin{array}{l} a_0(U, V) = \frac{1}{\rho v} (H_0 \tau, \nabla V)_{(L^2(\Omega))^4}, \\ a_1(U, V) = \frac{g}{C^2} (|U_0|U, V)_{(L^2(\Omega))^2}, \\ b_0(U, V) = (U_0 \cdot \nabla U, H_0 V)_{(L^2(\Omega))^2}, \\ b_1(U, V) = (lf]U, H_0 V)_{(L^2(\Omega))^2}. \end{array} \right. \tag{37}$$

where v is equal to μ_H/ρ in (8). It is noted that $\tilde{\tau}_{ij}$ in (8) is simply written as τ in Appendix A. Then we have

$$\begin{aligned} a(U, V) &= va_0(U, V) + a_1(U, V), \\ b(U, V) &= b_0(U, V) + b_1(U, V), \\ &\text{for } U, V \in (H_0^1(\Omega))^2. \end{aligned}$$

It is easy to see that there are positive constants $\alpha, \mu, \beta_0, \beta_1$ and γ with which the following estimates from (38) to (42) hold true:

$$a_1(V, V) \leq \alpha \|V\|_X^2 \quad \text{for } V \in (L^2(\Omega))^2. \tag{38}$$

$$a_1(V, V) \leq \mu a_0(V, V) \quad \text{for } V \in (H_0^1(\Omega))^2. \tag{39}$$

$$b_0(U, V) \leq \beta_0 a_0(U, U)^{1/2} \|V\|_X \quad \text{for } U \in (H_0^1(\Omega))^2, V \in (L^2(\Omega))^2, \tag{40}$$

$$\|V\|_X \leq \beta_1 a_0(V, V)^{1/2} \quad \text{for } V \in (H_0^1(\Omega))^2. \tag{41}$$

$$(\text{div}(H_0 U), \eta)_{L^2(\Omega)} \leq \gamma a_0(U, U)^{1/2} \|\eta\|_Z \quad \text{for } U \in (H_0^1(\Omega))^2, \eta \in L^2(\Omega). \tag{42}$$

In fact, (19) implies (38). Then (19), (3.12) in [15] (Korn’s first inequality) and (38) imply the estimates from (39) to (42). Since our family of “triangulation” $\{\mathfrak{T}_h : 0 <$

$h \leq \bar{h} < \infty$ satisfies the inverse assumption, there is a constant σ_0 independent of h satisfying

$$\|\nabla \zeta_h\|_{(L^2(\Omega))^2} \leq \frac{\sigma_0}{h} \|\zeta_h\|_{L^2(\Omega)} \quad \text{for } \zeta_h \in W_h.$$

Hence, there is the constant σ independent of h satisfying

$$a_0(V_h, V_h) \leq \left(\frac{\sigma}{h}\right)^2 \|V_h\|_X^2 \quad \text{for } V_h \in \mathcal{V}_h, \tag{43}$$

Theorem 4 Fix the parameter θ such that $0 \leq \theta \leq 1$. Let ε, δ , and $\tilde{\varepsilon}$ be positive constants satisfying

$$\varepsilon + \delta \leq 1$$

and

$$\tilde{\varepsilon} < 1,$$

respectively. For any $h \in (0, \bar{h}]$, choose τ so as to satisfy the condition:

$$\tau \leq 2\nu \min\left(\frac{\varepsilon}{(\sigma\sqrt{\nu^2 + \mu\nu + \theta\beta h})^2} h^2, \frac{\delta}{\gamma^2}\right), \tag{44}$$

and the condition:

$$\tau \leq \frac{(1 - \tilde{\varepsilon})\sqrt{\nu}}{(1 - \theta)\beta\sigma} h \tag{45}$$

if $\theta < 1$, where $\beta = \beta_0 + |f|\beta_1$. Let $\mathbf{u}_h(t)$ be the step function of t constructed by the solution \mathbf{u}_n of (E_{*h,τ*}) through the relation:

$$\mathbf{u}_h(t) = \mathbf{u}_n : n\tau \leq t < (n + 1)\tau, \quad n = 0, 1, 2, \dots$$

Denote $\mathbf{u}_h(0)$ by \mathbf{a}_h . Let $\mathbf{u}(t)$ be the generalized solution of (E). Under these settings, we have that

$$\|\mathbf{u}_h(t)\|_X \leq \|\mathbf{a}_h\|_X, \quad t \geq 0, \tag{46}$$

and that if

$$\lim_{h \rightarrow 0} \|\mathbf{a}_h - \mathbf{a}\|_X = 0, \tag{47}$$

then

$$\lim_{h \rightarrow 0} \|\mathbf{u}_h(t) - \mathbf{u}(t)\|_X = 0, \tag{48}$$

uniformly in $t \in [0, T]$ for any finite T .

Remark 2 Finally we remark that the function $\mathbf{u}_h(t)$ in Theorem 4 satisfies the error estimate:

$$\|\mathbf{u}_h(t) - \mathbf{u}(t)\|_X \leq Ch \quad \text{for any } t \in [0, T] \quad (49)$$

with an h -independent constant C provided that the solution $\mathbf{u}(t)$ is sufficiently smooth, and that the initial value \mathbf{a}_h is chosen so as to satisfy $\|\mathbf{a}_h - \mathbf{a}\|_X = O(h)$.

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