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STRATIGRAPHY OF PALEOGENE SEQUENCES IN WESTON - SIPITANG, SABAH

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ARTICLE DETAILS	ABSTRACT
<i>Article History:</i> Received 12 November 2017 Accepted 12 December 2017 Available online 1 January 2018	This study is aimed to renew the stratigraphy in Weston - Sipitang especially for Temburong Formation and Crocker Formation which are known as the Paleogene sequences. The interfingering relationship between Temburong Formation and Crocker Formation shows that the Temburong Formation is older than the Crocker Formation in terms of stratigraphic sequences which had been the opposite. Facies analysis, facies association studies and the correlation of lithostratigraphic units in the study area and earlier studies helped to interpret stratigraphy and renew the position of stratigraphic for the Paleogene sequences. Temburong Formation and Crocker Formation are dated as Upper Eocene to Lower Miocene age.

KEYWORDS

Temburong, Paleogene, Upper Eocene, Lower Miocene age, lithostratigraphic.

1. INTRODUCTION

The area of study comprising Weston to Sipitang is located in the southwest of Sabah (Figure 1) with latitude of $05^0 15'$ N to $05^0 01'$ N and longitude of $115^0 30'$ E to $115^0 39'$ E. Area of study mapped in geological map (Figure 1) is 482.08 km². This study focuses on Paleogene sequences consisting of Temburong Formation and Crocker Formation. Temburong Formation locality is located in Kg. Naluyan Besar (Tu), Kg. Sungai Tiga (T1, T2) and Kg. Ulu Sipitang (T/M) while Crocker Formation locality is located in Kg. Naparan (Cu, C1) and at Weston (C2).



Figure 1: Geological map of study area and Sabah map [1].

2. GENERAL GEOLOGY

The older part of Setap Shale Formation experienced the same folding deformation with the West Crocker Formation and later named Temburong Formation [2]. Previous studies show that the Crocker Formation is a part of the Temburong Formation based on the similarity of the rock units found in both of these formations [3]. Both of these formations have interfingering relationship to one another with the same stratigraphic units ranging from Eocene to Miocene identified by the presence of foraminifera fossils [4]. In Sipitang area comprising the West Crocker Formation and the Temburong Formation. The Crocker Formation is a turbidite sequence widely exposed along the west coast of Sabah with its major division classified as greywacke [5].

3. RESEARCH METHOD

Stratigraphic studies in the study area were focused on the study of lithostratigraphy through facies analysis and facies association studies. The method of lithostratigraphic unit's correlation was conducted by using lithology to determine the lateral distribution of facies and the change of facies. Stratigraphic studies were also described further in terms of stereotype and type locality, rock unit's relationship, fossil and the age of formation. The stratigraphic sequence model was constructed to describe the stratigraphic relationships between formations and also to explain the age of formations in study area.

4. FACIES ANALYSIS AND FACIES ASSOCIATION

The description of facies analysis and facies association can be found in Table 1.

Table 1: Facies analysis and facies association.

Rock Unit	Facies Analysis			Facies Association			Deposition
	Bed	Description	Interpretation	Туре	Description	Interpretation	Environment
Crocker Formation	Ta – Te (channel)	Bouma Ta is the bed marker with massive structure. The thickness of > 1 meter is described as the	It is deposited with high density of turbidite current. The grains are deposited at once as the current power decreases to form a massive lawer	Channel - levee	Channel is marked with the present of bouma Ta. Massive structure and levee which is an overflow of the channel	Channel-levee can be found in middle fan of deep marine and is affected by high density turbidite current and active sediment denocition	
		channel.	iorin a massive layer.		characterized by the presence of a	seument deposition.	

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					Td parallel lamination.		Middle fan (Deep marine)
	Ta/Tb – Te (channel)	Bouma Ta is the bed marker with massive structure and the thickness of > 1 meter. The layer begin with Ta or Tb.	It is deposited with high density of turbidite current. The grains are deposited at once as the current power decreases to form a massive layer.	Channel - lobe	Channel is marked with the present of bouma Ta and lobe is marked with the present of bouma Tb as the remains from the channel.	Channel-lobe can be found in middle fan of deep marine in a state of high energy flow regime.	
	Ta/Tb – Te (lobe)	Bouma Tb which has a parallel lamination structure is a lobe indicator and a layer that begins with Ta or Tb.	It is deposited in a state of higher flow regime with high energy to transport sand grains with traction.	Channel – lobe (sediment migration)	Channel is marked with the present of bouma Ta and lobe (sediment migration) is marked with the present of bouma Tc with cross	Lobe (sediment migration) is deposited on the edge of the middle fan after undergoing a transition of sediment deposition with low energy of	
	Tb- Te (lobe)	Bouma Tb with the parallel lamination structure is the lobe indicator.	It is deposited in a state of higher flow regime with high energy.		lamination structure. Channel is deposited on the lobe (sediment migration).	lower flow regime. Channel is deposited on top of it with high energy.	
Rock Unit	Facies Analysis		• • • •	Facies Association			Deposition
Crocker Formation	Td – Te (levee)	Bouma Td with parallel lamination structure is the indicator of levee	It is deposited with very little turbidite current flow. A slight change in the current causes	туре	Lobe is marked with the present of bouma Tb with parallel lamination	Lobe is deposited in the middle fan of deep marine with high energy flow as a	Lava onnen
		which is the overflow from channel.	lamination deposition of coarse and fine grains alternately.		structure. Lobe (sediment migration) is marked with the present of bouma	part of the sediment from the channel. Lobe (sediment migration) is deposited on the	Middle fan (deep marine)
	Tc – Te (lobe (sediment migration))	which is the overflow from channel. Bouma Tc with cross lamination structure undergone sediment migration.	lamination deposition of coarse and fine grains alternately. It is deposited in lower flow regime with enough energy to transport fine sand.	Lobe – lobe (sediment migration)	structure. Lobe (sediment migration) is marked with the present of bouma Tc with cross lamination structure and also a part of the sediment from the lobe.	part of the sediment from the channel. Lobe (sediment migration) is deposited on the edge of the middle fan after undergoing a transition of sediment deposition with low energy of lower flow regime.	Middle fan (deep marine)
	Tc – Te (lobe (sediment migration)) Tc –Te (migration lobe)	which is the overflow from channel. Bouma Tc with cross lamination structure undergone sediment migration. Bouma Tc with cross lamination structure.	lamination deposition of coarse and fine grains alternately. It is deposited in lower flow regime with enough energy to transport fine sand. It is deposited in lower flow regime with enough energy to transport fine sand.	Lobe – lobe (sediment migration) Migration lobe	structure. Lobe (sediment migration) is marked with the present of bouma Tc with cross lamination structure and also a part of the sediment from the lobe. Migration lobe is marked with the present of bouma Tc with cross lamination structure.	part of the sediment from the channel. Lobe (sediment migration) is deposited on the edge of the middle fan after undergoing a transition of sediment deposition with low energy of lower flow regime. Migration lobe is deposited in the outer fan of deep marine with lower flow regime with very little current flow.	Middle fan (deep marine) Outer fan (deep marine)

Rock Unit	Facies Analysis			Facies Association			Deposition
	Bed	Description	Interpretation	Туре	Description	Interpretation	Environment
Temburong	Te (basin	Bouma Te has no	It is deposited with		Basin plain is	It is deposited in the	Basin plain
Formation	plain)	structure and	very little current	Basin plain	marked with the	basin plain with very	
		sequence.	flow.		present of bouma	little current flow.	
					Te without any		
					sediment		
					structure.		

5. CORRELATION OF LITHOSTRATIGRAPHIC UNITS AND DEPOSITION ENVIRONMENT

5.1 Temburong Formation



Figure 2: Correlation of Temburong Formation and deposition environment.

Description:



5.2 Crocker Formation



Figure 3: Correlation of Crocker Formation and deposition environment.

Description:



6. THE RELATIONSHIP BETWEEN TEMBURONG FORMATION AND CROCKER FORMATION CAUSED BY REGRESSION PROCESS



Figure 4: The cycle sequence deposition of Temburong Formation and Crocker Formation.

7. STRATIGRAPHY

The description of stratotype, rock unit relationship, fossil and age are essential to further explain the stratigraphy in the study area.

7.1 Stratotype and Type Locality

For the Temburong Formation, there is no record for stratotype. However, the type locality of the Temburong Formation lies at the head of the river of Temburong river in Brunei [2]. The formation originates from Brunei Darussalam. According to a study, the stratotype for the Crocker Formation is located at Tenom Gorge [6]. There is no type locality record for the Crocker Formation. The origin of the Crocker Formation is located in the Crocker range at the western part of Sabah [7].

7.2 The Relationship of Rock Units

Temburong formation with the age of Upper Eocene to Lower Miocene is associated with six types of formations, which are the Crocker Formation, Meligan Formation, Setap Shale Formation, Mulu Formation, Kelalan Formation and Tubau Formation. In the study area, the interfingering relationship between Temburong Formation and Crocker Formation cannot be found. It can be found in Tenom which is also located in the southwestern part of Sabah [8]. There are nine formations that have a relationship with the Crocker Formation which has the same age as the Temburong Formation. Among them are the Trusmadi Formation, Pinosuk Gravel Formation, Chert-spilit Formation, Bongaya Formation, Kudat Formation, Temburong Formation, Meligan Formation, Belait Formation and Labang Formation.

7.3 Fossil and Age

The research on biostratigraphy for microfossil cannot be carried out because there is no trace of any microfossil in microfossil analysis. Therefore, only the lithostratigraphy study was concentrated. Due to this situation, no microfossil is observed in the study area. However, microfossil for rock units can be explained by reference to previous studies. Only trace fossil can be observed in the study area. For the Temburong Formation, no trace fossil can be found in the study area. However, there is a previous study record for the microfossil of Temburong Formation in the Tenom district which is also located in southwest of Sabah. There is a planktic foraminifera representing two biozones named Globorotalia ciperoensis zone (Upper Oligocene) and Catapsydrax dissimilis zone - Praeorbulina sicana (Lower Miocene) which are in the range of Upper Eocene to Lower Miocene age [9]. For the Crocker Formation, the trace fossil of *Helminthopsis* sp. can be found which proves that the Crocker Formation is deposited in the deep marine environment. There is a previous study record for the planktonic foraminifera of Crocker Formation named Globigerinoides sp. in the Tenom area which is also located in southwest of Sabah with the same age as the Temburong Formation [8].

8. STRATIGRAPHIC SEQUENCE MODEL



Figure 5: Stratigraphic sequence model of Weston – Sipitang, Sabah [4]

9. CONCLUSION

The interfingering relationship between Temburong Formation and Crocker Formation concluded that Temburong Formation is older than Crocker Formation in terms of stratigraphic sequence which had been the opposite. Both of these formations are dated as Upper Eocene to Lower Miocene.

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