



Environmental Processes

Subscriber access provided by UNIV OF CAMBRIDGE

Thicker shells compensate extensive dissolution in brachiopods under future ocean acidification

Emma Cross, Elizabeth M. Harper, and Lloyd S. Peck

Environ. Sci. Technol., Just Accepted Manuscript • DOI: 10.1021/acs.est.9b00714 • Publication Date (Web): 29 Mar 2019 Downloaded from http://pubs.acs.org on April 3, 2019

Just Accepted

"Just Accepted" manuscripts have been peer-reviewed and accepted for publication. They are posted online prior to technical editing, formatting for publication and author proofing. The American Chemical Society provides "Just Accepted" as a service to the research community to expedite the dissemination of scientific material as soon as possible after acceptance. "Just Accepted" manuscripts appear in full in PDF format accompanied by an HTML abstract. "Just Accepted" manuscripts have been fully peer reviewed, but should not be considered the official version of record. They are citable by the Digital Object Identifier (DOI®). "Just Accepted" is an optional service offered to authors. Therefore, the "Just Accepted" Web site may not include all articles that will be published in the journal. After a manuscript is technically edited and formatted, it will be removed from the "Just Accepted" Web site and published as an ASAP article. Note that technical editing may introduce minor changes to the manuscript text and/or graphics which could affect content, and all legal disclaimers and ethical guidelines that apply to the journal pertain. ACS cannot be held responsible for errors or consequences arising from the use of information contained in these "Just Accepted" manuscripts.



is published by the American Chemical Society. 1155 Sixteenth Street N.W., Washington, DC 20036

Published by American Chemical Society. Copyright © American Chemical Society. However, no copyright claim is made to original U.S. Government works, or works produced by employees of any Commonwealth realm Crown government in the course of their duties.

1	Thicker shells compensate extensive dissolution in
2	brachiopods under future ocean acidification
3	Emma L. $Cross^{\dagger, \ \ddagger, \ \ast}$, Elizabeth M. Harper [†] , Lloyd S. Peck [‡]
4	
5	⁺ Department of Earth Sciences, University of Cambridge, Downing Street,
6	Cambridge, CB2 3EQ, UK.
7	[‡] British Antarctic Survey, Natural Environment Research Council, High Cross,
8	Madingley Road, Cambridge, CB3 0ET, UK
9	
10	KEYWORDS: climate change, global warming, shell dissolution, shell thickness,
11	compensatory mechanisms, phenotypic plasticity, Calloria inconspicua, Liothyrella uva,
12	Terebratulide

14 ABSTRACT

15 Organisms with long generation times require phenotypic plasticity to 16 17 survive in changing environments 18 until genetic adaptation can be 19 achieved. Marine calcifiers are



20 particularly vulnerable to ocean acidification due to dissolution and a reduction in 21 shell-building carbonate ions. Long-term experiments assess organisms' abilities to 22 acclimatise or even adapt to environmental change. Here we present an unexpected 23 compensatory response to extensive shell dissolution in a highly calcium-carbonate-24 dependent organism after long-term culture in predicted end-century acidification 25 and warming conditions. Substantial shell dissolution with decreasing pH posed a 26 threat to both a polar (Liothyrella uva) and a temperate (Calloria inconspicua) 27 brachiopod after 7 months and 3 months exposure, respectively, with more extensive 28 dissolution in the polar species. This impact was reflected in decreased outer 29 primary layer thickness in the polar brachiopod. A compensatory response of 30 increasing inner secondary layer thickness, and thereby producing a thicker shell 31 was exhibited by the polar species. Less extensive dissolution in the temperate 32 brachiopod did not affect shell thickness. Increased temperature did not impact shell 33 dissolution or thickness. Brachiopod ability to produce a thicker shell when extensive shell dissolution occurs suggests this marine calcifier has great plasticity in 34

- 35 calcification providing insights into how similar species might cope under future
- 36 environmental change.

38 INTRODUCTION

39 Changing environments pose serious risks to organisms that cannot shift their geographic range, physiologically acclimatise or genetically adapt¹. Current 40 41 understanding of the biological impacts of ocean acidification and warming is 42 largely based on short- (days) to medium-term (weeks) laboratory and field experiments that have revealed mixed responses in many species²⁻⁵. More recently, 43 44 however, there has been an increase in long-term (many months to years) studies that demonstrate surprising capacities of marine organisms to acclimate⁶⁻¹⁰, or even 45 adapt in organisms with short generation times¹¹⁻¹³ to decreased pH and increased 46 47 temperature. Compensatory mechanisms could be paramount to maintain overall 48 performance of organisms that have limited capacities to alter their geographic range 49 under future changed conditions and subsequently sustain their key ecological 50 functions in our oceans¹⁴.

Marine calcifiers are considered the most vulnerable organisms to ocean 51 52 acidification due to the combination of dissolution and the reduction in carbonate 53 ions making shell production more difficult and energetically expensive^{2,15,16}. The 54 Southern Ocean has naturally low carbonate ion saturation levels compared to temperate and tropical regions due to carbon dioxide being more soluble in cold 55 56 water¹⁷. Acid-base coefficients are also more sensitive in cold temperatures making 57 this high latitude region a forerunner of biological ocean acidification impacts for 58 other oceans¹⁸. The external skeleton is crucial for protecting animal tissue in shell-59 bearing organisms against predation, infection and loss of bodily fluids^{19,20}. Any 60 negative impacts to shell integrity, therefore, could compromise its protective 61 function and potentially prove fatal. Shell integrity may be affected by erosion from 62 natural scour or attack from shell-boring organisms as well as dissolution. The 63 calcified shell of all shell-bearing organisms is protected by an outer organic layer, 64 the periostracum^{21,22}. Abrasion of this protective layer and subsequently inner shell 65 layers naturally occurs through abrasion from suspended inorganic particulate material, the movement of individuals against each other, and with other calcified 66 67 biota or substrata. Shell dissolution also poses a threat depending on the solubility of 68 the biomineral, the chemical characteristics of the surrounding seawater and 69 metabolic by-products released by the adhering biofilm^{23,24}. Predicted environmental 70 conditions for 2100 will shift surface seawater carbonate chemistry to favour CaCO₃ 71 dissolution, which could exacerbate the loss of shell integrity of marine calcifiers.

72 Compensatory mechanisms may counteract deleterious ocean acidification and 73 warming effects on organisms. For these to succeed, the compensatory mechanism 74 must occur at a faster rate than that of the deleterious effect to provide successful 75 protection. Phenotypic plasticity of shell morphology has been reported in shelled organisms in response to the presence of predators²⁵ and changing environmental 76 77 conditions^{14,26-28}. These include shell thickening, production of a more rotund shell 78 and increased shell growth rates through plasticity in producing different calcium 79 carbonate polymorphs^{14,25-32}. Production of a thicker periostracum could also withstand more wear and deter dissolution³³. Periostracum loss or shell dissolution 80 81 at the external surface far away from the secretory tissue cannot be directly repaired by the organism. Compensatory mechanisms such as induced thickening, however,could counteract this potentially fatal effect of ocean acidification.

84 Brachiopods are one of the most calcium-carbonate-dependent groups of marine 85 animals because their calcareous skeleton and other support structures make up > 86 90% of their dry mass^{34,35}. Rhynchonelliform brachiopods possess a low-magnesium calcite shell consisting of the periostracum underlain by two biomineralised inner 87 88 layers; the thin nanocrystalline primary layer and the generally much thicker fibrous 89 secondary layer^{36,37}. In previously published work we showed that shell growth 90 rates of L. uva Broderip, 1833 (which we refer to as "polar brachiopod") and C. 91 inconspicua Sowerby, 1846 (which we refer to as "temperate brachiopod") were not 92 impacted by predicted end-century seawater pH's^{6,7}. Another study demonstrated 93 increased dissolution in the polar brachiopod in pH 7.4 conditions after 14 days³⁸, 94 however, empty dried valves were used so the brachiopods ability to compensate 95 shell dissolution remains unknown. This study, therefore, investigated dissolution 96 effects and potential compensatory mechanisms of a polar and a temperate 97 brachiopod living under acidified and warming conditions. Specifically, the extent of 98 dissolution and thickness of whole valves and individual shell layers were assessed 99 under predicted end-century pH levels in both brachiopods and also under 100 increased temperature in the polar brachiopod.

101

102

MATERIALS AND METHODS

104

105 Sampling collection. Specimens of the polar brachiopod were hand collected by 106 SCUBA divers from Trolval Island, Ryder Bay, Antarctica (67° 35.44′ S, 68° 12.44′ W) 107 at 15-25 m depth in May 2012. Environmental conditions in Ryder Bay at 15-25 m depth consist of seawater temperatures that range from -1.8 to +1.5°C, however, 108 109 temperatures rarely exceed +1.0°C and salinity is 33.0-34.0³⁹ and the pH range is 110 8.04-8.10⁴⁰. Brachiopods were kept in recirculating aquaria ($0.0 \pm 0.5^{\circ}$ C) whilst being 111 transported by ship back to the British Antarctic Survey, Cambridge, UK where the 112 polar experiment was conducted. 113 Individuals of the temperate brachiopod were hand collected at low tide from 114 under rocks in Portobello Bay, Otago Harbour, New Zealand (45° 82.00'S, 170° 115 70.00'E) in January 2013. Environmental conditions in Otago Harbour are surface 116 seawater temperatures of 6.4-16.0°C^{41,42}, pH range of 8.10-8.21 (K. Currie, pers. 117 comm.) and salinity is 32.5-34.8⁴². Brachiopods were kept in seawater during the 118 short transportation to Portobello Marine Laboratory, Otago Harbour, New Zealand 119 where the temperate experiment was performed.

120

121 Experimental Design.

122 *Polar experiment.* The polar experiment was conducted in a temperature-123 controlled recirculating CO_2 microcosm with four treatments⁶. Two were acidified 124 treatments ("Moderate pH" – pH 7.75 ± 0.03 and "Low pH" – pH 7.54 ± 0.03) based 125 on the IPCC 'business-as-usual' scenario of the predicted end-century reduction of 126 0.3-0.5 pH units from the present day average of pH 8.1 in surface oceanic seawater by 2100⁴³ (Table 1). The third was a pH control where the seawater remained at 127 128 ambient pH (pH 8.05 ± 0.03). All these three treatments were maintained at 2°C 129 throughout the experiment due to the concurrent 2°C increase in sea surface 130 temperature (SST) expected to occur alongside these predicted decreased pH levels 131 by the end of the century⁴⁴. The fourth treatment was a temperature control which 132 was held at the present-day average conditions for Ryder Bay⁴⁵ (SST: 0°C, pH: 7.98 ± 133 0.02). The pH of the acidified treatments was controlled by intermittently bubbling 134 CO₂ gas into a header tank. Seawater was then gravity fed into the experimental 135 tanks⁶. The pH control treatment had a similar set up but without the pH 136 manipulation system. The temperature control treatment was situated separately in 137 the main BAS aquarium. Seawater temperature of all treatments was manipulated 138 by controlling the air temperature in temperature-controlled laboratories.

139 Seawater temperatures (°C, Digital Testo 106) and pH_{NIST} (Aquamedic pH 140 controlled computer and electrode system) were monitored and recorded daily. 141 Salinity (Tropical Marine Centre V2 Handheld refractometer), TCO₂ (mmol L⁻¹; Ciba 142 Corning TCO₂ Analyzer 965, Olympic Analytical. UK) and nutrient content (silicate 143 and phosphate) of each treatment were measured weekly. Other carbonate system 144 parameters, including the partial pressure of CO_2 (pCO_2) and the saturation values 145 for calcite ($\Omega_{\rm C}$) and aragonite ($\Omega_{\rm A}$), were modelled from applying TCO₂ and pH_{NIST} data to the program CO2SYS⁴⁶ with refitted constants^{47,48}. Brachiopods in each 146

treatment were fed weekly with microalgal concentrate of approximately 331 x 10^4 cells L⁻¹, which is within the natural seasonal range of phytoplankton cell abundance along the west Antarctic Peninsula (62–1150 x 10^4 cells L⁻¹) 49,50 .

150 Temperate experiment. The temperate experiment was conducted in a flow-151 through CO₂ perturbation system with three treatments⁷. Two were acidified 152 treatments ("Moderate pH" – pH 7.79 \pm 0.06 and "Low pH" – pH 7.62 \pm 0.05) and the 153 third was a pH control (8.16 ± 0.03). The pH of the acidified treatments was lowered 154 in header tanks by intermittently bubbling CO₂ gas before being gravity fed into the 155 replicate experimental tanks⁷. The pH control system had an identical set up except 156 that it lacked CO₂ injection, and air was injected into the header tank. Seawater 157 temperature was not manipulated and was ambient for Otago Harbour.

158 Seawater temperatures (°C, Digital Testo 106) and pH_{NIST} were measured three 159 times a day and salinity (YSI data logger) was measured once a week. Dissolved 160 inorganic carbon (DIC) and total alkalinity (A_T) were analysed at the beginning, middle and end of the experiment by a Single Operator Multi-parameter Metabolic 161 162 Analyser (SOMMA) and closed-cell potentiometric titration, respectively⁵¹. Other 163 carbonate system parameters, including the partial pressure of CO_2 (pCO_2) and the 164 saturation values for calcite ($\Omega_{\rm C}$) and aragonite ($\Omega_{\rm A}$) were calculated using CO2SYS⁴⁶ 165 with CO₂ equilibrium constants^{47,48,52}. Brachiopods were fed three times a week with microalgal concentrate of approximately 397 x 10⁴ cells mL⁻¹ of Tetraselmis spp., 166 which is within the natural summer range of phytoplankton cell abundance in Otago 167 168 Harbour.

169 **Table 1.** Mean (±SD) seawater parameters during both the polar and temperate

170 experiments.

Experiment	Treatment	pH _{NIST}	Temperature (°C)	Salinity	pCO ₂ (µatm)	Ω Calcite	Ω Aragonite
	Temperature control	7.98 ± 0.02	-0.3 ± 0.1	35 ± 1	417 ± 15	1.2 ± 0.1	0.8 ± 0.1
Polar	pH control	8.05 ± 0.03	1.7 ± 0.3	35 ± 1	365 ± 67	1.5 ± 0.2	0.9 ± 0.1
	Moderate pH	7.75 ± 0.03	1.9 ± 0.4	35 ± 1	725 ± 133	0.8 ± 0.1	0.5 ± 0.1
	Low pH	7.54 ± 0.03	2.2 ± 0.4	35 ± 1	1221 ± 179	0.5 ± 0.1	0.3 ± 0.1
	pH control	8.16 ± 0.03	16.5 ± 1.7	34 ± 1	465 ± 83	3.5 ± 0.5	2.2 ± 0.3
Temperate	Moderate pH	7.79 ± 0.06	16.9 ± 1.7	34 ± 1	1130 ± 12	1.6 ± 0.0	1.0 ± 0.0
	Low pH	7.62 ± 0.05	16.6 ± 1.7	34 ± 1	1536 ± 235	1.3 ± 0.2	0.8 ± 0.1

171 Values for pCO_2 , Ω calcite and Ω aragonite were calculated from CO2SYS⁴⁶ with 172 refitted constants^{47,48}.

173

174 Shell condition index. Shell lengths were measured at the start and end of each 175 experiment using Vernier calipers (± 0.1 mm) to determine shell laid down in the 176 natural environment that thickens from the internal surface as brachiopods grow 177 (which we refer to as "thickening shell") and shell growth extension during the 178 experiments (which we refer to as "growing shell"). Mean lengths (± S.E.) of these 179 two shell regions from each treatment are reported in Table S1. Scanning Electron 180 Microscopes (JEOL 820 for the polar brachiopod and FEI QEMSCAN 650F for the 181 temperate brachiopod; both operated using an accelerating voltage of 20 kV) were 182 used to image gold-coated outer surfaces of five ventral valves of adult specimens 183 from each treatment of both species to determine shell condition. Five types of shell 184 condition were present: intact shell (IS; intact periostracum with pitted layer), 185 minimal wear (W1; periostracum without pitted layer), extensive wear (W2; wear 186 but no dissolution), partial shell dissolution (SD1; dissolution in the inner primary 187 layer) and extensive shell dissolution (SD2; dissolution exposing the innermost 188 secondary layer). Full descriptions and examples of each type of shell condition for 189 both species are presented in Table S2. Micrographs (1 mm x 1 mm) were collected at 190 five standardised areas in thickening shell (areas located from umbo region towards 191 anterior margin as detailed in Fig. S1A) and five standardised areas in growing shell 192 (areas evenly spread in anterior margin as detailed in Fig. S1A). Percentage areas of 193 each type of shell condition from each SEM micrograph were calculated/measured in 194 ImageJ (Fig. S1B). Each shell region was analysed separately to determine whether 195 treatment and/or the location of shell analysed (which we refer to as "shell position") 196 affected shell that had already been potentially subjected to substantial wear 197 (thickening shell) and newly produced shell with less time subjected to wear 198 (growing shell).

199 Shell thickness. Longitudinal cross sections of five dorsal values of adult 200 specimens from each treatment of both species were finely polished to 3 μ m using 201 Kemet met papers (P400, P800, P2500 and P4000) followed by MetPrep diamond 202 solutions (6 μ m and 3 μ m). Acetate peels from polished cross sections of the brachial 203 values of both species were made according to a previous study⁵³. Thickness 204 measurements (± 0.1 mm) of the primary layer, secondary layer and total shell were 205 then measured from three areas of thickening shell (umbo region, middle of the shell and nearer experimental growth as detailed in Fig. S2) and three areas of growing shell (oldest experimental growth to newest experimental growth in the anterior margin as detailed in Fig. S2) on a Swift monocular petrological microscope with fitted micrometer.

210 Statistical analyses. Shell condition index data were non-normally distributed due to the presence of zeros in the dataset. Non-parametric Kruskal-Wallis tests 211 212 were, therefore, used to determine whether treatment and/or shell position affected 213 the median percentage area of each type of shell condition. When significant 214 differences occurred, post-hoc Dunn's tests were conducted to identify which 215 treatments and shell positions were statistically different from each other. As shell 216 condition and shell thickness measurements were conducted at several points within 217 an individual, Kruskal-Wallis tests were also used to determine if individual number 218 affected each shell condition. Linear mixed effects models were computed to 219 determine if treatment, shell position (fixed effects) and/or individual number 220 (random effect) impacted primary layer, secondary layer and total shell thickness:

221

222 Thickness measurement = Treatment + Shell Position + (1|Individual Number)

223

+ error

224

Likelihood ratio tests were used to determine p values (p < 0.05) between the full model with the effect in question against the reduced model without the effect in question. When the ratio tests identified significant differences, post-hoc Tukey tests

were performed to determine which treatments or shell positions were responsible. 228 229 Shell thickness data were checked for variance homogeneity and normality using 230 Levene's and Shapiro-Wilk tests (p < 0.05), respectively. Each shell region was analysed separately for both shell condition index and shell thickness to determine 231 232 whether treatment and/or shell position affected shell maintenance (thickening shell) and shell production (growing shell). Statistical analyses were computed using R⁵⁴ 233 with the FSA package⁵⁵ used for the Kruskal-Wallis and post-hoc Dunn's tests, the 234 *lme4* package⁵⁶ for the linear mixed effects models and the *emmeans* package⁵⁷ for the 235 236 post-hoc Tukey tests.

237

238 RESULTS

239

240 Shell condition index.

241 Thickening shell. Intact shell (IS) was absent from both acidified treatments and 242 only present in $< 8.4 \pm 4.5\%$ (mean \pm SE) of both controls in the thickening shell in the 243 polar brachiopod (Figure 1a & Figure 2). Instead, minimal wear (W1) dominated this 244 region in both controls (Figure 2a; 71.0 \pm 6.2% in pH control and 65.1 \pm 4.1% in 245 temperature control). Partial shell dissolution (SD1), however, was the most 246 prominent shell condition in both acidified treatments (Figure 1a; 64.3 ± 4.6% in 247 moderate pH and 71.7 ± 4.3% in low pH). With decreasing pH, the percentage area 248 of partial shell dissolution increased (Figure 1a & Figure 2; Kruskal-Wallis: H = 70.93, 249 p < 0.001). The extent of shell dissolution in the polar brachiopod also increased with 250 decreasing pH (Kruskal-Wallis: H = 42.38, p < 0.001), with 18.2 \pm 4.5% of shell 251 exhibiting exposed secondary layer in the low pH treatment (SD2) compared to 0.9 ± 252 0.4% in the moderate pH treatment and the secondary layer never being exposed in 253 either control. Temperature had no effect on shell dissolution or wear (Figure 1a; 254 Dunn's Test: SD1 – T = 1.16, p = 0.244, SD2 – T = 0.26, p = 0.795, W1 – T = 0.25, p = 255 0.805, W2 – T = 1.60, p = 0.109). In contrast to the polar brachiopod, thickening shell 256 of the temperate brachiopod was mainly characterised by intact shell (Figure 1c; IS; 257 56.6 - 82.3%) across all treatments. Amounts of minimal wear (W1) decreased with 258 decreasing pH in this shell region in the temperate brachiopod (Figure 1c; Kruskal-259 Wallis: H = 7.92, p = 0.020). Partial shell dissolution (SD1), however, increased with decreasing pH (Figure 1c & Figure 3; Kruskal-Wallis: H = 53.72, p < 0.001) in growing shell in the temperate brachiopod. Shell dissolution in this temperate species was less extensive than for the polar species (Figure 2 & Figure 3) as the secondary layer was not exposed (SD2) in any individual in any treatment. Shell position or individual number did not affect any shell condition in the thickening shell of both species (Table S3).

266 Growing shell. Growing shell in both species was mainly characterised by intact shell (IS) in all treatments (Figure 1b, d; polar brachiopod: > $63.9 \pm 4.7\%$; temperate 267 268 brachiopod: $> 83.2 \pm 1.8\%$). Less intact shell occurred in the most acidified conditions 269 compared to all other treatments in both species (Figure 1b, d; Kruskal-Wallis: polar 270 brachiopod - H = 41.81, p < 0.001; temperate brachiopod – H = 20.96, p < 0.001). 271 Partial shell dissolution (SD1) increased with increasing acidity in the experimental 272 growth of the polar brachiopod (Figure 1b & Figure 2; Kruskal-Wallis: polar brachiopod - H = 63.08, p < 0.001). This shell dissolution, however, occurred at a 273 274 much lower level ($3.2 \pm 1.0\%$ in moderate pH and $35.9 \pm 4.7\%$ in low pH) in the 275 growing shell than in the thickening shell in this species. Temperature had no effect 276 on partial shell dissolution (Dunn's Test: Temperature control vs pH control: T = -277 0.22, p = 0.829). Partial shell dissolution (SD1) only occurred in the most acidified 278 treatment in the temperate brachiopod (Figure 1d; $11.1 \pm 1.5\%$), also in lower levels 279 than in the thickening shell $(28.3 \pm 3.2\%)$ in this species. Extensive shell dissolution 280 (SD2) was absent in the growing shell in both species across all treatments. Minimal 281 wear (W1) was only present in two individuals across all treatments in the polar

brachiopod and was absent from the temperate brachiopod. Extensive wear (W2) was present in higher levels in the control treatments of both species than in the acidified treatments (Figure 1c, d; Kruskal-Wallis; polar brachiopod – H = 43.98, p < 0.001, temperate brachiopod - H = 10.67, p < 0.001), however, only in low levels (< 11.1 \pm 1.2%). Neither shell position nor individual number affected any shell condition in the growing shell of both species (Table S3).

288



Polar brachiopod



ACS Paragon Plus Environment



Figure 2. Representative shell condition in the polar brachiopod – Examples of SEM
micrographs of shell surfaces of thickening shell (A, C, E, G) and growing shell (B, D,
F, H) in temperature control (A, B), pH control (C, D), moderate pH (E, F) and low
pH treatment (G, H). IS = intact shell, W1 = minimal wear, W2 = extensive wear, SD1
= partial shell dissolution and SD2 = extensive shell dissolution. Scale bar = 20 µm.



303

Figure 3. Representative shell condition in the temperate brachiopod – Examples of SEM micrographs of shell surfaces of thickening shell (A, C, E) and growing shell (B, D, F) in pH control (A, B), moderate pH (C, D) and low pH treatment (E, F). IS = intact shell, W1 = minimal wear, W2 = extensive wear and SD1 = partial shell dissolution. SD2 (extensive shell dissolution) was absent in all treatment in this species. Scale bar = $20 \mu m$.

311 Shell thickness.

312 Thickening shell. The outer primary layer became progressively thinner in thickening shell as pH reduced in the polar brachiopod (Figure 4a; Linear Mixed 313 Effects Model; χ^2 = 79.72, df = 3, p < 0.001). Secondary layer thickened in all 314 315 treatments as this inner shell layer transitioned from the region of growing shell to 316 thickening shell as the brachiopod grew (larger grey bars in Figure 4a vs smaller 317 grey bars in Figure 4b). This inner secondary layer, and the whole shell, however, 318 were thicker in the most acidified treatment in the thickening shell (Figure 4a; Linear Mixed Effects Model; Secondary Layer - χ^2 = 39.63, df = 3, p < 0.001; Total Shell - χ^2 = 319 320 18.19, df = 3, p < 0.001). Increased temperature had no effect on primary layer, secondary layer or total shell thickness (Tukey; Primary Layer – T = 1.73, p = 0.319; 321 322 Secondary Layer – T = -1.20, p = 0.627; Total Shell – T = -0.80, p = 0.855). In contrast, 323 neither individual shell layers nor total shell thickness were affected by lowered pH 324 in the thickening shell in the temperate brachiopod (Figure 4c; Linear Mixed Effects Model; Primary Layer - χ^2 = 4.17, df = 2, p = 0.124; Secondary Layer - χ^2 = 4.80, df = 2, 325 p = 0.091; Total Shell - χ^2 = 4.27, df = 2, p = 0.118). Primary layer was thinnest in the 326 327 oldest part of the shell, the umbo region, across all treatments in both species (Table S4). Secondary layer and total shell thickness did not differ in different places in the 328 329 thickening shell in each treatment in both species (Table S4). Individual number also 330 had no effect on individual shell layer and total shell thickness in both species (Table 331 S4).

332 Growing shell. Primary layer, secondary layer and total shell thickness were not impacted by lowered pH in either species (Figure 4b & 4d; Linear Mixed Effects 333 Model; polar brachiopod: Primary Layer - χ^2 = 3.62, df = 3, p = 0.306; Secondary Layer 334 - χ^2 = 6.80, df = 3, p = 0.078; Total Shell - χ^2 = 5.26, df = 3, p = 0.154; temperate 335 brachiopod: Primary Layer - χ^2 = 2.63, df = 2, p = 0.268; Secondary Layer - χ^2 = 0.82, df 336 = 2, p = 0.663; Total Shell - χ^2 = 1.12, df = 3, p = 0.572). Increased temperature also had 337 338 no effect on either individual shell layers or total shell thickness in the polar 339 brachiopod (Figure 4b). Primary layer thickness did not differ indifferent places 340 throughout the growing shell in each treatment in either species (Table S4). 341 Secondary layer and the total shell thickness, however, did get progressively thinner 342 with the direction of growth in each treatment in both species (Table S4). Individual 343 number also had no effect on individual shell layer and total shell thickness in 344 growing shell in either species (Table S4).



Polar brachiopod



Figure 4. Shell thickness – Mean primary layer (black bar) and secondary layer (grey bar) thicknesses from three areas in the thickening shell (a, b) and from three areas in the growing shell (c, d) in the polar brachiopod (top row, n = 5 per treatment) and in the temperate brachiopod (bottom row, n = 5 per treatment) in all treatments. Whole bars represent total shell thickness. Lowercase letters a-f indicate significant differences in primary layer thickness and g-k represent significant differences in secondary layer and total shell thicknesses between treatments in each shell region

in each species. Comparisons were made only within shell region not between shellregions or between species.

355

356 DISCUSSION

Long-term culturing of a polar and a temperate brachiopod under predicted endcentury acidified conditions revealed that both species were more susceptible to shell dissolution with increasing acidity. Our two principal findings are significant dissolution and an unexpected compensation of induced thicker shells in the thickening shell.

Dissolution of shell. Shell loss has been widely reported in several marine 362 calcifiers, however, these have largely been those which use higher solubility 363 364 polymorphs of calcium carbonate (i.e. aragonite), such as corals⁵⁸⁻⁶⁰ and molluscs^{23,32,61-64}, high-magnesium calcite including coralline 365 algae^{65,66} and 366 echinoderms^{67,68}. Fewer studies have investigated shell dissolution in taxa which are 367 entirely constructed of the lower solubility polymorph, low-magnesium calcite, such as rhynchonelliform brachiopods. Previously, the only other ocean acidification 368 369 study assessing dissolution in brachiopods was conducted on dead shells³⁸. Working 370 on the polar species, they showed deterioration of the primary layer after only 35 371 days exposure to pH 7.4, which after 56 days exposed the fibres of the secondary 372 layer below. This is the same dissolution pattern reported here in experiments 373 involving live individuals after 7 months exposure to pH 7.54. Exposure of the

374 secondary layer calcite fibres may compromise shell integrity and probably strength
375 due to the loss of the hard outer protective primary layer^{38,69}.

376 Dissolution was more extensive in the polar than in the temperate brachiopod, as 377 indicated by increased deterioration in the primary layer of the polar species 378 compared to the temperate brachiopod in the moderate pH treatment and it was 379 only in the polar species that the secondary layer was exposed in the low pH 380 treatment. Antarctic calcified invertebrates are probably the most vulnerable 381 organisms to ocean acidification for a number of reasons: they tend to be weakly 382 calcified^{16,70}; dissolution rates of calcium carbonate are inversely related to 383 temperature¹⁷; and the polar regions are predicted to become the first to be 384 undersaturated in aragonite by 2050 and calcite by 2100^{18,40,71-74}. Both the moderate 385 pH and the low pH treatment in the polar experiment were undersaturated with 386 respect to calcite, however, both the acidified treatments in the temperate 387 experiment were not undersaturated with respect to calcite. This could explain the 388 differences in the extent of dissolution present between both investigated species. 389 The state of the shells could have also influenced these species differences. Wear was 390 more prominent in the thickening shell of the polar brachiopod than in the 391 temperate brachiopod, which was most likely due to the longer lifespan of the polar 392 species (up to 55-60 years)⁷⁵ compared to the temperate species (up to 14 years)⁷⁶. 393 Thus the shells of the polar brachiopod had been exposed to wear for a longer time 394 in their natural environment before the experiment began. Such wear will have 395 damaged or removed periostracum, which is key in protecting the animal from shell 396 dissolution^{31,77-82}. Since periostracum is only formed at the growing edge of the 397 mantle, it cannot be repaired if damaged or lost from the surface of the shell. 398 Thinning or loss of this organic layer through physical or biotic abrasion and 399 epibiont erosion, therefore, restricts protection from corrosive acidified waters. The 400 periostracum in brachiopods is < 1 μ m thick⁸³ and so is very vulnerable to loss.

Newly formed growing shell was mainly characterised by intact shell in both 401 402 species. Partial shell dissolution did occur, however, in the most acidified treatment 403 in both species albeit at a much lower level than in the thickening shell. Damage to 404 the ultrathin periostracum from abrasion of other brachiopods in their conspecific 405 cluster, natural decay of this outer layer or potentially the lowered pH conditions 406 could have either softened the periostracum itself or disrupted the protective 407 function of the periostracum. This latter possibility was suggested for external 408 dissolution reported in newly formed shell in M. edulis after 2 months exposure to 1400 µatm and 4000 µatm⁷⁹. Disintegration of organic matrix in the shell rather than 409 410 corrosion of crystals could have caused this shell degradation, as seen in spirorbirds 411 after 100-day exposure to pH 7.7 conditions⁸⁴.

Temperature had no clear effect on shell dissolution or thickness in the polar brachiopod as indicated by the lack of or only minimal primary layer dissolution and no change in any thickness measurement in both thickening and growing shell in the temperature control (held at 0° C – current average Antarctic summer temperatures) and the pH control (kept at the 2°C temperature increase predicted for 2100). In contrast, temperature and not acidification reduced shell strength in *M*. 418 edulis after 6 months exposure to forecasted end-century pH and warming 419 conditions⁸⁵. It was concluded that warming had an indirect effect on shell strength 420 by shifting the energy budget from shell deposition to increased maintenance costs. 421 Food availability was limited throughout the experiment, which would likely have 422 enhanced the temperature effect as low food levels can reduce shell growth and 423 significantly influence the amount of inner shell dissolution in *M. edulis* after 7 weeks exposure to varying pCO_2 levels⁸⁶. This highlights the necessity of using 424 425 multistressors in ocean acidification research to better understand the abilities of 426 marine calcifiers to maintain shell integrity under future predicted environmental 427 conditions.

Compensation. Despite the widely reported significant effects of dissolution on 428 marine calcifiers in ocean acidification research, very few studies investigate 429 430 organisms' abilities to compensate for shell loss. New shell deposited by M. edulis 431 after 9 months exposure to 750 μ atm and 1000 μ atm pCO₂ was rounder and flatter 432 with a thinner aragonite layer than shell produced in ambient conditions of 380 433 µatm²⁷. The authors attributed this new shell shape to a compensatory mechanism to 434 enhance protection from predators and changing environments as these mussels 435 were unable to grow thicker shells in high pCO_2 conditions. Shell thickening has 436 occurred in response to biotic shell loss by endoliths and other conspecifics grazing 437 on their external shell in Patellid limpets, Patella granatina and P. argenvillei⁸⁷, and to 438 abiotic shell loss by physical impacts from ice in the Antarctic limpet Nacella 439 concinna⁸⁸. Decreased shell thickness has also been reported in molluscs in lowered

pH conditions, due to internal dissolution of the highly soluble aragonite layer^{27,78,86}. 440 441 For compensatory mechanisms to succeed, they must occur at faster rates than the 442 deleterious effect. Thicker basal shells were reported in the barnacle Amphibalanus 443 amphitrite under lowered pH conditions (pH 7.4), however, this compensation 444 calcification was insufficient as dissolution weakened shells faster than it was 445 deposited⁸⁹. A pteropod specimen collected from the Fram Strait in the Arctic Ocean 446 also produced a shell four times thicker than the original shell in response to mechanical and dissolution damage from undersaturated waters³¹. 447

Extensive shell dissolution at low pH in thickening shell of the polar brachiopod 448 449 led to a drastic decrease in primary layer thickness. The polar species counteracted 450 this chemical attack by laying down more secondary layer on the internal surface of the shell, which resulted in increased overall shell thickness during the experimental 451 452 period. The less extensive dissolution in the temperate brachiopod was reflected by 453 no clear impact of acidified conditions on either total shell or individual shell layer 454 thicknesses. Our findings appear to contrast with reports of primary layer thickening 455 in the Chilean terebratulide Magellania venosa after being cultured in pH 7.35 456 conditions⁹⁰, however, their observations appear to be based on only one measured 457 specimen in both the acidified treatment and the control.

Compensatory mechanisms must also be sufficient in maintaining an organism's overall performance. The secondary layer of terebratulide brachiopod shells is softer than the harder protective primary layer^{37,91} raising the question of whether a shell made solely out of secondary layer would provide adequate protection to ensure survival. No external dissolution of the exposed secondary layer of the polar brachiopod was observed perhaps due to protection from the organic matrix shrouding calcite crystals of this innermost fibrous shell layer^{92,93}. Primary layer is often missing in older parts of brachiopod shells or in older individuals⁹⁰, therefore, a thicker shell consisting of only secondary layer could provide sufficient protection in predicted pH conditions expected by 2100. Although, ocean acidification impacts on brachiopod shell strength warrant further investigation.

469 Total shell thickness or individual shell layer thickness of growing shell of both 470 species were not affected by predicted end-century acidified conditions. Shell 471 thickness, therefore, is only impacted by lowered pH when extensive shell 472 dissolution occurs. In a previous study, shell thickness in the temperate brachiopod 473 did not vary over the last 120 years despite a 0.1 pH unit decrease and 2°C increase 474 in temperature since the Industrial Revolution⁹⁴. Forecasted acidified conditions by 475 2100 also did not impact shell growth rates and the ability to shell repair in both the 476 polar and temperate brachiopod^{6,7}. The resilience of shell thickness in both the polar 477 and temperate species to past and predicted environmental change, in addition to 478 their unaffected shell growth rates under end-century pH levels^{6,7}, indicates the 479 robust ability of rhynchonelliform brachiopods to construct shell under acidified 480 conditions. The thickness of calcite and aragonite layers in newly formed shell of *M*. edulis were also not affected by elevated pCO_2^{79} . This lack of variation in shell 481 482 thickness to acidified conditions in newly produced shell further demonstrates the 483 increase of shell thickness in the thickening shell is a compensatory response to

extensive shell dissolution occurring at the external shell surface, although the mechanisms whereby the brachiopods identify the shell is thinning remain to be elucidated.

487 The extent of vulnerability of two highly calcium-carbonate-dependent species to 488 dissolution in acidified seawater is concerning. Without any counteracting response, 489 dissolution may compromise shell integrity leading to reduced protection and 490 decreased suitability of brachiopod shells as a habitat for other marine organisms. 491 Physiological acclimatisation is one approach organisms can utilise to cope with 492 such threats in the challenging conditions predicted by 2100. We identified induced 493 shell thickening forming thicker shells in the polar brachiopod as a compensatory 494 mechanism to extensive shell dissolution under lowered pH levels. The less 495 extensive dissolution in the temperate species was probably a function of higher 496 temperatures in the temperate study and the corresponding lower CaCO₃ solubility. 497 This suggests that the level of dissolution in the temperate brachiopod after 3 498 months exposure to predicted end-century pH conditions did not induce similar 499 compensation. This induced shell thickening could come at an overall cost to the 500 organism as increased shell production is energy-demanding, involving the 501 accumulation, transportation and precipitation of calcium carbonate as well as the 502 production of the organic matrix^{95,96}. Acidification also significantly increases the 503 proportion of the animal's energy budget that needs to be devoted to shell 504 production⁹⁷, therefore, there may be long-term impacts on life histories and 505 maintenance of populations. Long-term experiments investigating the capacity of organisms to acclimatise and possibly adapt to future change is crucial to further ourunderstanding of how marine organisms will cope with future climate change.

Marine organisms may also adjust physiological, behavioural or ecological traits as 508 509 additional compensatory responses to their changing habitats. As well as direct 510 effects on energy budgets (e.g. induced shell thickening), ocean acidification could 511 also have indirect impacts through the alteration of their resource quality (e.g. 512 energy intake)98. To maintain organismal homeostasis in varying environments, 513 individuals may compensate by modifying the quality and quantity of food 514 consumed, which in turn could also stabilise community productivity⁹⁹. Multiple 515 compensatory mechanisms could be paramount to maintain overall performance of organisms and subsequently sustain key community processes under future 516 517 environmental change.

518

519 ASSOCIATED CONTENT

520 Supporting Information

521 Mean lengths (\pm S.E) of thickening and growing shell regions (Table S1)

522 Descriptions and examples of each shell condition (Table S2)

523 Schematic and example of shell condition index measurements (Figure S1)

- 524 Schematic and example of shell thickness measurements (Figure S2)
- 525 Shell dissolution additional statistical results (Table S3)
- 526 Shell thickness additional statistical results (Table S4)

527	
528	AUTHOR INFORMATION
529	Corresponding Author
530	* Emma L. Cross. Email: E.L.Cross@cantab.net. Address: Department of Earth
531	Sciences, University of Cambridge, Downing Street, Cambridge, CB2 3EQ, UK and
532	British Antarctic Survey, Natural Environment Research Council, High Cross,
533	Madingley Road, Cambridge, CB3 0ET, UK
534	ACKNOWLEDGEMENTS
535	The authors would like to thank the scientific dive team at the British Antarctic
536	Survey Rothera Research Station for collecting the L. uva specimens and to Dr.
537	Coleen Suckling, Dr. Simon Morley and Rebecca Smith for their help in the set-up
538	and maintenance of the polar experimental system. We are very grateful to Dr. Miles
539	Lamare for his assistance in the organisation and collection of the C. inconspicua
540	specimens. Thanks also to the science support staff at the Portobello Marine
541	Laboratory, University of Otago for their help in the construction and maintenance
542	of the temperate experimental system. Thanks to Dr. Kim Currie at the National
543	Institute of Water and Atmospheric Research for the DIC and total alkalinity
544	measurements. This research was funded by the NERC PhD Studentship
545	(NE/T/A/2011) awarded to ELC.

547

548 REFERENCES

- 549 (1) Peck, L. S., Organisms and responses to environmental change. *Mar.*550 *Genom.* 2011, 4, 237-243.
- (2) Doney, S. C.; Fabry, V. J.; Feely, R. A.; Kleypas, J. A., Ocean acidification:
 the other CO₂ problem. *Ann. Rev. Mar. Sci.* 2009, *1*, 169-192.
- (3) Hofmann, G. E.; Barry, J. P.; Edmunds, P. J.; Gates, R. D.; Hutchins, D. A.;
 Klinger, T.; Sewell, M. A., The effect of ocean acidification on calcifying organisms in
 marine ecosystems: An organism-to-ecosystem perspective. *Annu Rev Ecol Evol Syst*2010, *41*, 127-147.
- (4) Byrne, M., Impact of ocean warming and ocean acidification on marine
 invertebrate life history stages: vulnerabilities and potential for persistence in a
 changing ocean. *Oceanogr Mar Biol Annu Rev* 2011, 49, 1-42.
- (5) Parker, L. M.; Ross, P. M.; O'Connor, W. A.; Pörtner, H. O.; Scanes, E.;
 Wright, J. M., Predicting the response of molluscs to the impact of ocean
 acidification. *Biology* 2013, *2*, 651-92.
- (6) Cross, E. L.; Peck, L. S.; Harper, E. M., Ocean acidification does not impact
 shell growth or repair of the Antarctic brachiopod *Liothyrella uva* (Broderip, 1833). *J. Exp. Mar. Biol. Ecol.* 2015, 462, 29-35.
- (7) Cross, E. L.; Peck, L. S.; Lamare, M. D.; Harper, E. M., No ocean
 acidification effects on shell growth and repair in the New Zealand brachiopod *Calloria inconspicua* (Sowerby, 1846). *ICES J. Mar. Sci.* 2016, 73, 920-926.
- (8) Queirós, A. M.; Fernandes, J. A.; Faulwetter, S.; Nunes, J.; Rastrick, S. P.;
 Mieszkowska, N.; Artioli, Y.; Yool, A.; Calosi, P.; Arvanitidis, C.; Findlay, H. S.;
 Barange, M.; Cheung, W. W.; Widdicombe, S., Scaling up experimental ocean
 acidification and warming research: from individuals to the ecosystem. *Global Change Biol.* 2015, *21*, 130-143.
- 574 (9) Suckling, C. C.; Clark, M. S.; Richard, J.; Morley, S. A.; Thorne, M. A.; 575 Harper, E. M.; Peck, L. S., Adult acclimation to combined temperature and pH

576 stressors significantly enhances reproductive outcomes compared to short-term
577 exposures. J. Anim. Ecol. 2014, 84, 773-784.

(10) Hazan, Y.; Wangensteen, O. S.; Fine, M., Tough as a rock-boring urchin:
adult *Echinometra* sp. EE from the Red Sea show high resistance to ocean
acidification over long-term exposures. *Mar. Biol.* 2014, *161*, 2531-2545.

(11) Collins, S.; Rost, B.; Rynearson, T. A., Evolutionary potential of marine
phytoplankton under ocean acidification. *Evol. Appl.* 2014, *7*, 140-155.

(12) Kelly, M. W.; Hofmann, G. E., Adaptation and the physiology of ocean
acidification. *Funct. Ecol.* 2013, 27, 980-990.

(13) Sunday, J. M.; Calosi, P.; Dupont, S.; Munday, P. L.; Stillman, J. H.;
Reusch, T. B., Evolution in an acidifying ocean. *Trends Ecol. Evol.* 2014, 29, 117-125.

(14) Leung, J. Y. S.; Russell, B. D.; Connell, S. D., Mineralogical plasticity acts
as a compensatory mechanism to the impacts of ocean acidification. *Environ. Sci. Technol.* 2017, *51*, 2652-2659.

590 (15) Byrne, M.; Przeslawski, R., Multistressor impacts of warming and
591 acidification of the ocean on marine invertebrates' life histories. *Integ. Comp. Biol.*592 2013, 53, 582-596.

(16) Watson, S.-A.; Peck, L. S.; Tyler, P. A.; Southgate, P. C.; Tan, K. S.; Day, R.
W.; Morley, S. A., Marine invertebrate skeleton size varies with latitude, temperature
and carbonate saturation: implications for global change and ocean acidification. *Global Change Biol.* 2012, *18*, 3026-3038.

597 (17) Revelle, R. R.; Fairbridge, R. W., Carbonates and carbon dioxide. In
598 *Geological Society of America Memoirs*, 1957; Vol. 67, pp 239-296.

(18) Fabry, V. J.; McClintock, J. B.; Mathis, J. T.; Grebmeier, J. M., Ocean
acidification at high latitudes: the bellwether. *Oceanography* 2009, 22, 160-171.

(19) Vermeij, G. J., The Mesozoic marine revolution: evidence from snails,
predators and grazers. *Paleobiology* 1977, *3*, 245-258.

(20) Harper, E. M.; Clark, M. S.; Hoffman, J. I.; Philipp, E. E.; Peck, L. S.;
Morley, S. A., Iceberg scour and shell damage in the Antarctic bivalve *Laternula elliptica*. *PLoS ONE* 2012, 7, e46341.

606 (21) Harper, E. M., The molluscan periostracum: an important constraint in
607 bivalve evolution. *Palaeonotology* **1997**, *40*, 71-97.

608 (22) Williams, A.; Mackay, S., Secretion and ultrastructure of the periostracum
609 of some terebratulide brachiopods. *Proc. R. Soc. B.* **1978**, 202, 191-209.

(23) Nienhuis, S.; Palmer, A. R.; Harley, C. D., Elevated CO₂ affects shell
dissolution rate but not calcification rate in a marine snail. *Proc. R. Soc. B.* 2010, 277,
2553-2558.

(24) Bausch, A. R.; Gallego, M. A.; Harianto, J.; Thibodeau, P.; Bednaršek, N.;
Havenhand, J. N.; Klinger, T., Influence of bacteria on shell dissolution in dead
gastropod larvae and adult *Limacina helicina* pteropods under ocean acidification
conditions. *Mar. Biol.* 2018, 165, 40.

- 617 (25) Freeman, A. S.; Byers, J. E., Divergent induced responses to an invasive
 618 predator in marine mussel populations. *Science* 2006, *313*, 831-834.
- (26) Peyer, S. M.; Hermanson, J. C.; Lee, C. E., Developmental plasticity of
 shell morphology of quagga mussels from shallow and deep-water habitats of the
 Great Lakes. *J. Exp. Biol.* 2010, *213*, 2602-2609.

(27) Fitzer, S. C.; Vittert, L.; Bowman, A.; Kamenos, N. A.; Phoenix, V. R.;
Cusack, M., Ocean acidification and temperature increase impact mussel shell shape
and thickness: problematic for protection? *Ecol. Evol.* 2015, *5*, 4875-4884.

- (28) Telesca, L.; Michalek, K.; Sanders, T.; Peck, L. S.; Thyrring, J.; Harper, E.
 M., Blue mussel shell shape plasticity and natural environments: a quantitative approach. *Sci. Rep.* 2018, *8*, 2865.
- 628 (29) Vermeij, G. J., Phenotypic evolution in a poorly dispersing snail after
 629 arrival of a predator. *Nature* 1982, 299, 349-350.
- (30) Fisher, J. A.; Rhile, E. C.; Liu, H.; Petraitis, P. S., An intertidal snail shows
 a dramatic size increase over the past century. *PNAS* 2009, *106*, 5209-12.

(31) Peck, V. L.; Oakes, R. L.; Harper, E. M.; Manno, C.; Tarling, G. A.,

633	Pteropods counter mechanical damage and dissolution through extensive shell
634	repair. Nat. Commun. 2018, 9, 264.
635	(32) Duquette, A.; McClintock, J. B.; Amsler, C. D.; Pérez-Huerta, A.; Milazzo,
636	M.; Hall-Spencer, J. M., Effects of ocean acidification on the shells of four
637	Mediterranean gastropod species near a CO ₂ seep. <i>Mar. Pollut. Bull.</i> 2017 .
638	(33) Telesca, L.; Peck, L. S.; Sanders, T.; Thyrring, J.; Sejr, M. K.; Harper, E. M.,
639	Plasticity and environmental heterogeneity predict geographic resilience patterns of
640	foundation species to future change. <i>bioRxiv</i> 2018 .
641	(34) Peck, L. S., The tissues of articulate brachiopods and their value to
642	predators. Phil. Trans. R. Soc. B. 1993, 339, 17-32.
643	(35) Peck, L. S., Brachiopods and climate change. Earth Env. Sci. T. R. So. 2008,
644	98, 451-456.
645	(36) Williams, A.; Brunton, C. H. C.; MacKinnon, D. I., Morphology. In Treatise
646	on Invertebrate Paleontology, Part H, Brachiopods (Revised), Kaesler, R. L., Ed. The
647	Geological Society of America and The University of Kansas Press: Boulder,
648	Colorado, and Lawrence, Kansas, 1997; Vol. 1, pp 321-422.
649	(37) Goetz, A. J.; Griesshaber, E.; Neuser, R. D.; Lüter, C.; Hühner, M.; Harper,
650	E. M.; Schmahl, W. W., Calcite morphology, texture and hardness in the distinct
651	layers of rhynchonelliform brachiopod shells. <i>Eur. J. Mineral.</i> 2009 , <i>21</i> , 303-315.
652	(38) McClintock, J. B.; Angus, R. A.; McDonald, M. R.; Amsler, C. D.;
653	Catledge, S. A.; Vohra, Y. K., Rapid dissolution of shells of weakly calcified Antarctic
654	benthic macroorganisms indicates high vulnerability to ocean acidification. Antarct.
655	<i>Sci.</i> 2009, <i>21,</i> 449-456.

(39) Venables, H. J.; Clarke, A.; Meredith, M. P., Wintertime controls on
summer stratification and productivity at the western Antarctic Peninsula. *Limnol. Oceanogr.* 2013, *58*, 1035-1047.

(40) McNeil, B. I.; Matear, R. J., Southern Ocean acidification: a tipping point
at 450-ppm atmospheric CO₂. *PNAS* 2008, *105*, 18860-18864.

(41) Greig, M. J.; Ridgway, N. M.; Shakespeare, B. S., Sea surface temperature
variations at coastal sites around New Zealand. *N. Z. J. Mar. Freshwat. Res.* 1988, 22,
391-400.

(42) Roper, D. S.; Jillett, J. B., Seasonal occurrence and distribution of flatfish
(Pisces: Pleuronectiformes) in inlets and shallow water along the Otago coast. N. Z. J. *Mar. Freshwat. Res.* 1981, 15, 1-13.

(43) IPCC, Climate Change 2013: The Physical Science Basis. In *Working Group I Contribution to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*, Stocker, T. F.; Qin, D.; Plattner, G.-K.; Tignor, M.; Allen, S. K.; Boschung, J.;
Nauels, A.; Xia, Y.; Bex, V.; Midgley, P. M., Eds. Cambridge, United Kingdom and
New York, NY, USA, 2013; p 1552.

(44) Mitchell, J. F. B.; Senior, C. A.; Johns, T. C., *Transient response to increasing greenhouse gases using models with and without flux adjustment*. Metrological Office:
1998.

(45) Clarke, A.; Meredith, M. P.; Wallace, M. I.; Brandon, M. A.; Thomas, D.
N., Seasonal and interannual variability in temperature, chlorophyll and
macronutrients in northern Marguerite Bay, Antarctica. *Deep Sea Res. Part II Top. Stud. Oceanogr.* 2008, 55, 1988-2006.

(46) Lewis, E.; Wallace, D. W. R.; Allison, L. J., *Program developed for CO₂ system calculations*. Carbon Dioxide Information Analysis Center, Oak Ridge National
Laboratory, US Department of Energy, Oak Ridge, Tennessee, US, 1998.

(47) Mehrbach, C.; Culberson, C. H.; Hawley, J. E.; Pytkowicxz, R. M.,
Measurement of apparent dissociation constants of carbonic acid in seawater at
atmospheric pressure. *Limnol. Oceanogr.* 1973, *18*, 897-907.

(48) Dickson, A. G.; Millero, F. J., A comparison of the equilibrium-constants
for the dissociation of carbonic-acid in seawater media. *Deep-Sea Res Pt I* 1987, 34,
1733-1743.

688	(49) Garibotti, I. A.; Vernet, M.; Ferrario, M. E., Annually recurrent
689	phytoplanktonic assemblages during summer in the seasonal ice zone west of the
690	Antarctic Peninsula (Southern Ocean). Deep-Sea Res Pt I 2005, 52, 1823-1841.
691	(50) Garibotti, I. A.; Vernet, M.; Ferrario, M. a. E.; Smith, R. C.; Ross, R. M.;
692	Quentin, L. B., Phytoplankton spatial distribution patterns along the western
693	Antarctic Peninsula (Southern Ocean). Mar. Ecol. Prog. Ser. 2003, 261, 21-39.
694	(51) Dickson, A. G.; Sabine, C. L.; Christian, J. R. Guide to best practices for
695	<i>Ocean</i> CO_2 <i>Measurements</i> ; 2007.
696	(52) Wanninkhof, R.; Lewis, E.; Feely, R. A.; Millero, F. J., The optimal
697	carbonate dissociation constants for determining surface water pCO_2 from alkalinity
698	and total inorganic carbon. Mar. Chem. 1999 , 65, 291-301.
699	(53) Richardson, C. A.; Crisp, D. J.; Runham, N. W., Tidally deposited growth
700	bands in the shell of the common cockle, <i>Cerastoderma edula</i> (L.). <i>Malacologia</i> 1979 , <i>18</i> ,
701	277-290.
702	(54) R, Core Team, R: A language and environment for statistical computing.
703	R Foundation for Statistical Computing, Vienna, Austria. URL https://www.R-
704	<u>project.org/</u> . 2017.
705	(55) Ogle, D. H., Fisheries Stock Analysis. 2018 , R Package version 0.8.20.
706	(56) Bates, D.; Maechler, M.; Bolker, B.; Walker, S., Fitting linear mixed-effects
707	models using lme4. J. Stat. Softw. 2015, 67, 1-48.
708	(57) Lenth, R., emmeans: Estimated Marginal Means, aka Least-Squares
709	Means. 2018 , R package version 1.2.3.
710	(58) Comeau, S.; Carpenter, R. C.; Lantz, C. A.; Edmunds, P. J., Ocean
711	acidification accelerates dissolution of experimental coral reef communities.
712	Biogeosciences 2015 , 12, 365-372.
713	(59) Silbiger, N. J.; Donahue, M. J., Secondary calcification and dissolution
714	respond differently to future ocean conditions. <i>Biogeosciences</i> 2015, 12, 567-578.
715	(60) Andersson, A. J.; Kuffner, I. B.; MacKenzie, F. T.; Jokiel, P. L.; Rodgers, K.
716	S.; Tan, A., Net loss of CaCO ₃ from a subtropical calcifying community due to

seawater acidification: mesocosm-scale experimental evidence. *Biogeosciences* 2009, 6,
1811-1823.

(61) Hall-Spencer, J. M.; Rodolfo-Metalpa, R.; Martin, S.; Ransome, E.; Fine,
M.; Turner, S. M.; Rowley, S. J.; Tedesco, D.; Buia, M. C., Volcanic carbon dioxide
vents show ecosystem effects of ocean acidification. *Nature* 2008, 454, 96-9.

(62) Milano, S.; Schöne, B. R.; Wang, S.; Müller, W. E., Impact of high *p*CO₂ on
shell structure of the bivalve *Cerastoderma edule*. *Mar. Environ. Res.* 2016, 119, 144-155.

(63) Bednaršek, N.; Tarling, G. A.; Bakker, D. C. E.; Fielding, S.; Feely, R. A.,
Dissolution dominating calcification process in polar pteropods close to the point of
aragonite undersaturation. *PLoS ONE* 2014, *9*, e109183.

(64) Harvey, B. P.; Agostini, S.; Wada, S.; Inaba, K.; Hall-Spencer, J. M.,
Dissolution: The achilles' heel of the triton shell in an acidifying ocean. *Front. Mar. Sci.* 2018, *5*, 371.

(65) Kamenos, N. A.; Burdett, H. L.; Aloisio, E.; Findlay, H. S.; Martin, S.;
Longbone, C.; Dunn, J.; Widdicombe, S.; Calosi, P., Coralline algal structure is more
sensitive to rate, rather than the magnitude, of ocean acidification. *Global Change Biol.*2013, 19, 3621-3628.

(66) Cornwall, C. E.; Boyd, P. W.; McGraw, C. M.; Hepburn, C. D.; Pilditch, C.
A.; Morris, J. N.; Smith, A. M.; Hurd, C. L., Diffusion boundary layers ameliorate the
negative effects of ocean acidification on the temperate coralline macroalga *Arthrocardia corymbosa. PLoS ONE* 2014, 9, e97235.

(67) Miles, H.; Widdicombe, S.; Spicer, J. I.; Hall-Spencer, J., Effects of
anthropogenic seawater acidification on acid-base balance in the sea urchin *Psammechinus miliaris. Mar. Pollut. Bull.* 2007, 54, 89-96.

(68) Dubois, P., The skeleton of postmetamorphic echinoderms in a changing
world. *Biol. Bull.* 2014, 226, 223-236.

(69) Scurr, D. J.; Eichhorn, S. J., Structure/property relationships in seashells.
In *Mechanical Properties of Bioinspired Materials* Viney, C.; Katti, K.; Ulm, F.-J.;

Hellmich, C., Eds. Materials Research Society Symposium Proceedings: Warrendale,
PA, 2005; pp 87–92.

747 (70) Nicol, D., Some characteristics of cold-water marine pelecypods. *J.*748 *Paleontol.* **1967**, *41*, 1330-1340.

(71) Caldeira, K.; Wickett, M. E., Ocean model predictions of chemistry
changes from carbon dioxide emissions to the atmosphere and ocean. *J. Geophys. Res.*2005, *110*, C09S04.

(72) Orr, J. C.; Fabry, V. J.; Aumont, O.; Bopp, L.; Doney, S. C.; Feely, R. A.;
Gnanadesikan, A.; Gruber, N.; Ishida, A.; Joos, F.; Key, R. M.; Lindsay, K.; MaierReimer, E.; Matear, R.; Monfray, P.; Mouchet, A.; Najjar, R. G.; Plattner, G. K.;
Rodgers, K. B.; Sabine, C. L.; Sarmiento, J. L.; Schlitzer, R.; Slater, R. D.; Totterdell, I.
J.; Weirig, M. F.; Yamanaka, Y.; Yool, A., Anthropogenic ocean acidification over the
twenty-first century and its impact on calcifying organisms. *Nature* 2005, *437*, 681-6.

(73) Fabry, V. J.; Seibel, B. A.; Feely, R. A.; Orr, J. C., Impacts of ocean
acidification on marine fauna and ecosystem processes. *ICES J. Mar. Sci.* 2008, 65,
414-432.

761 (74) Guinotte, J. M.; Fabry, V. J., Ocean acidification and its potential effects on
762 marine ecosystems. *Ann. N. Y. Acad. Sci.* 2008, 1134, 320-42.

763 (75) Peck, L. S.; Brey, T., Bomb signals in old Antarctic brachiopods. *Nature*764 **1996**, *380*, 207-208.

(76) Doherty, P. J., A demographic study of a subtidal population of the New
Zealand articulate brachiopod *Terebratella inconspicua*. *Mar. Biol.* **1979**, *52*, 331-342.

767 (77) Ries, J. B.; Cohen, A. L.; McCorkle, D. C., Marine calcifiers exhibit mixed
768 responses to CO₂-induced ocean acidification. *Geology* 2009, *37*, 1131-1134.

(78) Tunnicliffe, V.; Davies, K. T. A.; Butterfield, D. A.; Embley, R. W.; Rose, J.
M.; Chadwick Jr, W. W., Survival of mussels in extremely acidic waters on a submarine volcano. *Nat. Geosci.* 2009, *2*, 344-348.

(79) Thomsen, J.; Gutowska, M. A.; Saphörster, J.; Heinemann, A.;
Trübenbach, K.; Fietzke, J.; Hiebenthal, C.; Eisenhauer, A.; Körtzinger, A.; Wahl, M.;

Melzner, F., Calcifying invertebrates succeed in a naturally CO₂-rich coastal habitat
but are threatened by high levels of future acidification. *Biogeosciences* 2010, *7*, 38793891.

(80) Rodolfo-Metalpa, R.; Houlbreque, F.; Tambutte, E.; Boisson, F.; Baggini,
C.; Patti, F. P.; Jeffree, R.; Fine, M.; Foggo, A.; Gattuso, J.-P.; Hall-Spencer, J. M., Coral
and mollusc resistance to ocean acidification adversely affected by warming. *Nature Clim. Change* 2011, *1*, 308-312.

(81) Coleman, D. W.; Byrne, M.; Davis, A. R., Molluscs on acid: gastropod
shell repair and strength in acidifying oceans. *Mar. Ecol. Prog. Ser.* 2014, 509, 203-211.

(82) Peck, V. L.; Tarling, G. A.; Manno, C.; Harper, E. M.; Tynan, E., Outer
organic layer and internal repair mechanism protects pteropod *Limacina helicina* from
ocean acidification. *Deep Sea Res. Part II Top. Stud. Oceanogr.* 2016, 127, 41-52.

(83) Williams, A.; MacKay, S., Differentiation of the brachiopod periostracum. *Palaeontology* 1979, 22, 721-736.

(84) Peck, L. S.; Clark, M. S.; Power, D.; Reis, J.; Batista, F. M.; Harper, E. M.,
Acidification effects on biofouling communities: winners and losers. *Global Change Biol.* 2015, *21*, 1907-1913.

(85) Mackenzie, C. L.; Ormondroyd, G. A.; Curling, S. F.; Ball, R. J.; Whiteley,
N. M.; Malham, S. K., Ocean warming, more than acidification, reduces shell
strength in a commercial shellfish species during food limitation. *PLoS ONE* 2014, *9*,
e86764.

(86) Melzner, F.; Stange, P.; Trubenbach, K.; Thomsen, J.; Casties, I.; Panknin, U.; Gorb, S. N.; Gutowska, M. A., Food supply and seawater pCO_2 impact calcification and internal shell dissolution in the blue mussel *Mytilus edulis*. *PLoS ONE* **2011**, *6*, e24223.

(87) Day, E. G.; Branch, G. M.; Viljoen, C., How costly is molluscan shell
erosion? A comparison of two patellid limpets with contrasting shell structures. *J. Exp. Mar. Biol. Ecol.* 2000, 243, 185-208.

802 (88) Hoffman, J. I.; Peck, L. S.; Hillyard, G.; Zieritz, A.; Clark, M. S., No
803 evidence for genetic differentiation between Antarctic limpet *Nacella concinna*804 morphotypes. *Mar. Biol.* 2009, 157, 765-778.

805 (89) McDonald, M. R.; McClintock, J. B.; Amsler, C. D.; Rittschof, D.; Angus,
806 R. A.; Orihuela, B.; Lutostanski, K., Effects of ocean acidification over the life history
807 of the barnacle *Amphibalanus amphitrite*. *Mar. Ecol. Prog. Ser.* 2009, *385*, 179-187.

(90) Ye, F.; Jurikova, H.; Angiolini, L.; Brand, U.; Crippa, G.; Henkel, D.;
Laudien, J.; Hiebenthal, C.; Šmajgl, D., Variation in brachiopod microstructure and
isotope geochemistry under low-pH–ocean acidification conditions. *Biogeosciences*2019, *16*, 617-642.

(91) Griesshaber, E.; Schmahl, W. W.; Neuser, R.; Pettke, T.; Blum, M.;
Mutterlose, J.; Brand, U., Crystallographic texture and microstructure of
terebratulide brachiopod shell calcite: An optimized materials design with
hierarchical architecture. *Am. Mineral.* 2007, *92*, 722-734.

816 (92) Pérez-Huerta, A.; Dauphin, Y.; Cusack, M., Biogenic calcite granules--are
817 brachiopods different? *Micron* 2013, 44, 395-403.

818 (93) Harper, E. M., Are calcitic layers an effective adaptation against shell
819 dissolution in the Bivalvia? *J. Zool. Lond.* 2000, 251, 179-186.

(94) Cross, E. L.; Harper, E. M.; Peck, L. S., A 120-year record of resilience to
environmental change in brachiopods. *Global Change Biol.* 2018, 24, 2262-2271.

822 (95) Palmer, A. R., Calcification in marine molluscs: How costly is it? *PNAS*823 **1992**, *89*, 1379-1382.

824 (96) Spalding, C.; Finnegan, S.; Fischer, W. W., Energetic costs of calcification
825 under ocean acidification. *Global Biogeochem. Cycles* 2017.

(97) Watson, S.-A.; Morley, S. A.; Peck, L. S., Latitudinal trends in shell
production cost from the tropics to the poles. *Sci. Adv.* 2017, *3*, e1701362.

828 (98) Ghedini, G.; Connell, S. D., Organismal homeostasis buffers the effects of
829 abiotic change on community dynamics. *Ecology* 2016, *97*, 2671-2679.

- 830 (99) Connell, S. D.; Ghedini, G., Resisting regime-shifts: the stabilising effect of
- 831 compensatory processes. *Trends Ecol. Evol.* **2015**, *30*, 513-5.

- 833 DISCLOSURES
- 834 The authors declare no competing financial interest.