



1 **CO₂ effects on diatoms: A Synthesis of more than a decade of ocean**
2 **acidification experiments with natural communities**

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12

13 **Abstract**

14 Diatoms account for ~40% of marine primary production and are considered to be key
15 players in the biological carbon pump. Ocean acidification (OA) is expected to affect
16 diatoms primarily by changing the availability of CO₂ as a substrate for photosynthesis
17 or through altered ecological interactions within the marine food web. Yet, there is little
18 consensus how entire diatom communities will respond to increasing CO₂. To address
19 this question, we synthesized the literature from over a decade of OA-experiments with
20 natural diatom communities to uncover: 1) if and how bulk diatom communities respond
21 to elevated CO₂; 2) if shifts within the diatom communities could be expected and how
22 they are expressed with respect to taxonomic affiliation and size structure. We found that
23 diatom communities responded to high CO₂ in ~60 % of the experiments and in this case
24 more often positively (56 %) than negatively (32 %; 12 % did not report the direction of



25 change). Shifts among different diatom species were observed in 65 % of the experiments.
26 Our synthesis supports the hypothesis that high CO₂ particularly favors larger species as
27 12 out of 13 experiments which investigated cell size found a shift towards larger species.
28 Unraveling winners and losers with respect to taxonomic affiliation was difficult due to
29 a limited database, but there is evidence that the genus *Pseudo-nitzschia* could be among
30 the losers. We conclude that OA-induced changes in diatom competitiveness and
31 assemblage structure must be classified as a “risk for ecosystem services” due to the
32 pivotal role diatoms play in trophic transfer and biogeochemical cycles.

33 **1. Introduction**

34 The global net primary production (NPP) of all terrestrial and marine autotrophs amounts
35 to approximately 105 petagrams (Pg) of carbon per year (Field et al., 1998). Marine
36 diatoms, a taxonomically diverse group of cosmopolitan phytoplankton, were estimated
37 to contribute up to 25 % (26 Pg C year⁻¹) to this number, which is more than the annual
38 primary production in any biome on land (Field et al., 1998; Nelson et al., 1995; Tréguer
39 and De La Rocha, 2013). Thus, diatoms are likely the most important single taxonomic
40 group of primary producers on Earth and any change in their prevalence relative to other
41 phytoplankton taxa could profoundly alter marine food web structures and thereby affect
42 ecosystem services such as fisheries or the sequestration of CO₂ in the deep ocean
43 (Armbrust, 2009; Tréguer et al., 2018).

44 The most conspicuous feature of diatoms is the formation of a silica shell, which is
45 believed to primarily serve as protection against grazers (Hamm and Smetacek, 2007;
46 Pančić and Kiørboe, 2018). Since the formation of this shell requires dissolved silicate,
47 diatoms are often limited by silicon as a nutrient rather than by nitrogen or phosphate
48 (Brzezinski and Nelson, 1996). However, when dissolved silicate is available, diatoms



49 benefit from their high nutrient uptake and growth rates, allowing them to outcompete
50 other phytoplankton and form intense blooms in many ocean regions (Sarhou et al.,
51 2005).

52 Diatoms display an enormous species richness, with recent estimates being in the range
53 of 30,000 species (Mann and Vanormelingen, 2013). Although only a fraction has been
54 morphologically described, known diatom taxa span a size range of several orders of
55 magnitude (<5 μm up to a few mm) with a wide range of morphologies and life strategies,
56 e.g. single cells and cell chains, pelagic and benthic habitats (Armbrust, 2009; Mann and
57 Vanormelingen, 2013; Sournia et al., 1991). Accordingly, they should not be treated as
58 one functional group, but rather as a variety of subgroups occupying different niches.

59 It is well recognized that the global importance of diatoms as well as their diversity in
60 morphology and life style is tightly linked to the functioning of pelagic food webs and
61 elemental cycling in the oceans. For example, iron enrichment experiments in the
62 Southern Ocean found that a shift in diatom community composition from thick- to thin-
63 shelled species (“persistence strategy” vs. “boom-and-bust strategy”) can enhance carbon
64 and alter nutrient export via sinking particles (Assmy et al., 2013; Smetacek et al., 2012).
65 This may not only affect element fluxes locally but enhance nutrient retention within the
66 Southern Ocean and reduce productivity in the north which underlines how important
67 diatom community shifts can be on a global scale (Boyd, 2013; Primeau et al., 2013;
68 Sarmiento et al., 2004). Likewise, the cell size of diatoms can play an important role in
69 transferring energy to higher trophic levels, as the dominance of larger species is
70 generally considered to reduce the length of the food chain and lead to higher trophic
71 transfer efficiency (Sommer et al., 2002). Consequently, understanding impacts of global
72 change on diatom community composition is crucial for assessing the sensitivity of
73 biogeochemical cycles and ecosystem services in the world oceans.



74 It has become evident that the sensitivity of diatoms to increasing pCO₂ is highly variable,
75 likely being related to specific traits such as cell size or the carbon fixation pathway, as
76 well as interactions with other environmental factors such as nutrient stress, temperature
77 or light (Gao et al., 2012; Hoppe et al., 2013; Wu et al., 2014). However, it is still rather
78 unclear how these species-specific differences in CO₂ sensitivities manifest themselves
79 on the level of diatom communities. This knowledge gap motivated us to compile the
80 presently available experimental data in order to reveal common responses of diatom
81 communities to high CO₂ and thereby assess potential scenarios of shifts in diatom
82 community composition under ocean acidification.

83 **2. Literature investigation**

84 **2.1. Approach**

85 Our original intention was to conduct a classical meta-analysis, which would have yielded
86 the benefit of a quantitative measure of diatom responses to OA, expressed as an overall
87 effect size (i.e. combined magnitude) such as the response ratio. However, our literature
88 analysis revealed a large variability in experimental pCO₂ ranges as well as measured
89 response variables, which cannot be directly compared among each other (e.g.
90 microscopic cell counts, pigment concentrations, genetic tools). These limitations impede
91 data aggregation as required for a classical meta-analysis. Furthermore, experimental
92 setups differed widely in terms of other environmental factors such as temperature, light,
93 and nutrient concentrations, all of which are known to modulate potential responses to
94 pCO₂ (Boyd et al., 2018), thereby further complicating data aggregation for meta-
95 analysis. Therefore, we chose an alternative, semi-quantitative approach where diatom
96 responses to increasing CO₂ are grouped in categories (see section 2.2) and also allows
97 to account for differences in experimental setups, e.g. with respect to container volume.



98 While this approach excludes the determination of effect size, it provides an unbiased
99 insight on the direction of change of potential CO₂ effects.

100 **2.2. Data compilation**

101 We explored the response of diatom assemblages to high CO₂ (low pH) by searching the
102 literature for relevant results with Google Scholar (December 15, 2017) using the
103 following search query: diatom OR Bacillariophyceae AND "ocean acidification" OR
104 "high CO₂" or "carbon dioxide" OR "elevated CO₂" OR "elevated carbon dioxide" OR
105 "low pH" OR "decreased pH". The first 200 results were inspected and considered to be
106 relevant when they were published in peer-reviewed journals, contained a description of
107 the relevant methodological details, a statistical analysis or at least a transparent
108 description of variance and uncertainties, and tested CO₂ effects on natural plankton
109 assemblages (artificially composed communities were not considered). We then carefully
110 checked the cited literature in these relevant studies to uncover other studies that were
111 missed by the initial search. Furthermore, we checked the "Ocean Acidification news
112 stream provided by the Ocean Acidification International Coordination Centre" under the
113 tag "phytoplankton" (<https://news-oceanacidification-icc.org/tag/phytoplankton/>) for
114 relevant updates since December 2017 (last check on January 16, 2019).

115 There were two response variables of interest for the literature compilation:

116 1) The response of the "bulk diatom community" to high CO₂. For this we checked if the
117 abundance of diatoms, the biomass of diatoms, or the relative portion of diatoms within
118 the overall phytoplankton assemblage increased or decreased under high CO₂ relative to
119 the control. We distinguished between "positive", "negative", and "no effect" following
120 the statistical results provided in the individual references. When the CO₂ effect on the
121 bulk community was derived from abundance data we also checked if there are



122 indications for a concomitant shift in the biomass distribution among species. This is
123 relevant because, for example, an increase in bulk abundance could coincide with a
124 decrease in bulk biomass when the species driving the abundances is smaller. We found
125 no indications for conflicting cases but acknowledge that not every reference provided
126 sufficient data on morphological details to fully exclude this scenario.

127 2) The CO₂-dependent species shifts within the diatom community with respect to
128 taxonomic composition and/or size structure. Unfortunately, cell size of the species was
129 not reported for all experiments. Thus, we distinguished between “no shifts”, “shifts
130 between species with unspecified size”, as well as “shifts towards larger or smaller
131 species” when this information was provided. Furthermore, we noted the winners and
132 losers within the diatom communities when these were reported (on the genus level).

133 In case the data was taken from factorial multiple stressor experiments (e.g. CO₂ x
134 temperature) we only considered the control treatment at ambient conditions (e.g. at
135 control temperature). Furthermore, we extracted various metadata from each study
136 largely following the literature analysis of (Schulz et al., 2017). All bulk diatom
137 responses, community shifts, and metadata is compiled/described in Table 1 and most of
138 it is self-explanatory (e.g. incubation temperature). The habitats of the investigated
139 diatom communities were categorized according to water depth, salinity, or life style in
140 the case of benthic communities: “oceanic” = water depth > 200 m (unless the habitat lies
141 within a fjord or fjord-like strait), S > 30; “coastal” = water depth < 200 m, S > 30;
142 “estuarine” = water depth < 200 m, S < 30; “benthic” = benthic communities (diatoms
143 growing on plates) were investigated. We reconstructed the water depth in case it was not
144 provided in the paper using Google Earth Pro (version 7.3.2.5495). The coordinates
145 provided in some of the experiments conducted in land-based facilities were imprecise
146 and marked positions on land. In this case the habitats were set to coastal or estuarine



147 depending on salinity. If salinity was not given we checked the location on Google Earth
148 for potential fresh water sources and also checked the text for more cryptic indications
149 (e.g. “euryhaline” in a lagoon were strong indications for an estuarine habitat). The
150 methods with which responses of the bulk diatom communities to high OA were
151 determined varied greatly among studies and included light microscopy (LM), pigment
152 analyses (PA), flow cytometry (FC), genetic tools (PCR), and biogenic silica (BSi)
153 analyses (Table 1).

154 **2.3. Balancing the influence of smaller and larger scale experiments to account for**
155 **the “degree of realism”**

156 The most realistic OA experiment would be one where all aspects of the natural habitat
157 are represented correctly. Such setups are possible for benthic communities which can be
158 sampled *in situ* along a natural CO₂ gradient at volcanic CO₂ seeps (Fabricius et al., 2011;
159 Hall-Spencer et al., 2008; Johnson et al., 2011). However, this does not work easily for
160 the large majority of pelagic studies compiled herein due to water advection. Thus, OA
161 experiments with pelagic communities are performed in closed containers which
162 inevitably cause artefacts (Calvo-Díaz et al., 2011; Ferguson et al., 1984; Guangao, 1990;
163 Menzel and Case, 1977). However, the degree by which they are unrealistic will differ
164 from study to study depending on the experimental design (Duarte et al., 1997). Here, we
165 aimed to develop a metric that allows us to estimate the realism of experiments with
166 pelagic communities in order to balance their influence on the final outcome of the
167 literature analysis. Most certainly, we do not mean to devalue any studies but think that
168 the highly different scales of experiments should not be ignored when evaluating the
169 literature. In the following we will first derive the equation for the proposed metric –
170 termed the “relative degree of realism (RDR)” – and introduce the underlying
171 assumptions. Afterwards we describe aspects that were considered while conceptualizing



172 the RDR.

173 The experimental design in the studies considered herein ranged from smaller bottle
174 experiments (e.g. 1 L) to *in situ* mesocosm studies with considerably larger incubation
175 volumes (e.g. 75000 L). While smaller differences in incubation volumes (e.g. 0.5 vs. 2
176 L) were shown to have no, or a minor, influence on physiological rates (Fogg and
177 Calvario-Martinez, 1989; Hammes et al., 2010; Nogueira et al., 2014; Robinson and
178 Williams, 2005), they can influence food web composition e.g. by excluding larger
179 grazers (Calvo-Díaz et al., 2011; Spencer and Warren, 1996). Larger differences of
180 incubation volumes (e.g. 10 vs. 10000 L) are considered to be important in all aspects
181 (Duarte et al., 1997), with the larger volume being more representative of natural
182 processes (Sarnelle, 1997). Therefore, our first assumption to conceptualize the RDR was
183 that larger incubation volumes represent nature generally better than smaller ones.

184 Plankton communities were pre-filtered in many experiments to exclude larger and often
185 patchily distributed organisms. This is a valid procedure to reduce noise and to increase
186 the likelihood to detect CO₂ effects but it also influences the development of plankton
187 communities as they modify the grazer/prey link within the food web (Ferguson et al.,
188 1984; Nogueira et al., 2014; Pomeroy et al., 1994). For example, (Nogueira et al., 2014)
189 compared plankton successions of pre-filtered (100 µm) and unfiltered communities and
190 found that the removal of larger grazers and diatoms gave room for green algae and
191 picophytoplankton to grow. Such manipulations make the experiment less representative
192 for a natural food web which brought us to the second assumption for the RDR: The
193 smaller the mesh size during the pre-filtration treatment, the less complete and thus the
194 less realistic is the pelagic food web.

195 To parameterize the two abovementioned assumptions we first converted the volume



196 information provided in each experiment into a volume-to-surface ratio (V/S). The
197 underlying thought is that V increases with the third power to the surface area of the
198 incubator and is indicative for the relation of open space to hard surfaces (Ferguson et al.,
199 1984). Therefore we first converted V into a radius (r) assuming spherical shape:

$$200 \quad r = \sqrt[3]{\frac{3V}{4\pi}} \quad (1).$$

201 The surface (S) of the spherical volume was calculated as:

$$202 \quad S = 4\pi r^2 \quad (2)$$

203 Assuming spherical shape was necessary because there is generally no information about
204 the shape of the incubation containers available. Although shape can influence processes
205 within the container (Pan et al., 2015), it is probably a less important factor to consider in
206 light of the large volume differences compared herein (Table 1).

207 The influence of pre-filtration treatments of the investigated plankton community is
208 implemented by multiplying the V/S with the third root of the applied mesh size (d_{mesh} in
209 μm) so that the RDR is defined as:

$$210 \quad \text{RDR} = \frac{V}{S} \sqrt[3]{d_{\text{mesh}}} \quad (3).$$

211 Thus, as for V/S, the influence of d_{mesh} on RDR does not linearly increase but dampens
212 with increasing d_{mesh} . The rationale for the non-linear increase is that incubations will still
213 have an increasing bias even if they do not have any pre-filtration treatment due to
214 generally increasing organism motility with size. For example, when collecting a
215 plankton community with a Niskin bottle, more motile organisms can escape from the
216 approaching sampler so that the food web composure is still affected even without



217 subsequent pre-filtration. For this reason we also capped the maximum d_{mesh} to 10,000
218 μm when there was no pre-filtration treatment applied since none of the studies included
219 significantly larger organisms. Figure 1 illustrates the change of RDR as a function of V
220 and d_{mesh} . High RDRs are calculated for large-scale *in situ* mesocosm studies ($\sim 50 - 190$)
221 while bottle experiments yield RDRs between $\sim 1 - 12$.

222 The key pre-requisite for an experimental parameter to be included in the RDR equation
223 (eq. 3) was that it is reported in all studies. Many parameters that we would have liked to
224 use for the RDR are either insufficiently reported (e.g. the light environment) or not
225 provided quantitatively at all (e.g. turbulence). We therefore had to work with very basic
226 properties related to the experimental setup rather than to the experimental conditions.

227 A particularly critical aspect of the RDR we had to deal with was the duration of the
228 experiments (Time). Time is a quantity, which is reliably reported in all studies and
229 therefore principally suitable for the RDR. Our first thoughts were that a realistic
230 community experiment should be long enough to cover relevant ecological processes
231 such as competitive exclusion and therefore also parameterized Time in the first versions
232 of the RDR equation. However, we decided to not account for it in the final version
233 because the factors that define the optimal duration of an experiment are poorly
234 constrained. For example, a 1 day experiment in a 10 L container could indeed miss
235 important CO_2 effects caused by food web interactions. On the other hand, a 30 days
236 experiment in the same container could reveal such indirect effects but at the same time
237 be associated with profound bottle effects and make the study unrepresentative for
238 simulated natural habitat. Thus, too long and too short are both problematic and the
239 optimum is hard to find. One such attempt to find the optimum was made by (Duarte et
240 al., 1997) who analyzed the plankton ecology literature between 1990 – 1995. By
241 correlating the experimental duration with the incubation volume of published



242 experiments they provided an optimal length for any given volume. However, as noted
243 by (Duarte et al., 1997), their correlation is based on publication success and therefore
244 rather reflects common practice in plankton ecology experiments and not necessarily a
245 mechanistic understanding of bottle effects. Thus, as there is no solid ground for a
246 parameterization of Time we ultimately decided to not consider it for the RDR.

247 **3. Results**

248 We found 54 relevant publications on CO₂ experiments with natural diatom assemblages.
249 Some publications included more than one experiment so that 69 experiments are
250 considered hereafter (Table 1). Most were done with diatom communities from coastal
251 environments (46 %) and oceanic (28%) environments. Estuarine and benthic
252 communities were investigated in 16 % and 6% of the studies, respectively (Figs. 2 and
253 3). 4 % of the studies did not provide coordinates where the samples were taken although
254 the region was reported (Table 1; Fig. 3).

255 One third (33 %) of all experiments revealed a positive influence of CO₂ on the “bulk
256 diatom community” (see section 2.2), while 19 % revealed a negative one. 7 % of the
257 studies found a CO₂ effect but did not specify whether it is a positive or negative one. 41
258 % found no effect (Fig. 4A; left column). Those experiments that revealed positive CO₂
259 effects on bulk diatom communities yielded the highest cumulative RDR score (\sum RDR)
260 of 605 while the \sum RDR for negative CO₂ effects was 266. No CO₂ effects yielded a score
261 of 768 while an “unspecified effect” yielded 266.

262 CO₂-dependent shifts in diatom species composition were investigated with light
263 microscopy except for (Endo et al., 2015) who used molecular tools,. Species shifts were
264 investigated in a subset of 40 of the 69 experiments (Fig. 4B). Within this subset of 40
265 studies, 12 (30 %, \sum RDR = 265) found a shift towards larger diatom species under high



266 CO₂, 1 (2.5 %, Σ RDR = 10) found a shift towards smaller diatom species, and 13 (32.5
267 %, Σ RDR = 103) found no CO₂ effect on diatom community composition. 14 studies (35
268 %, Σ RDR = 141) reported a CO₂-dependent shift but did not further specify any changes
269 in the size-class distribution (Fig. 4C).

270 A taxon-specific assessment of potential winners and losers (on the genus level) was
271 possible only to a limited extent, because most genera were not present in enough
272 experiments to get useful results. Only *Chaetoceros*, *Cylindrotheca*, and *Pseudo-*
273 *Nitzschia* were explicitly investigated in at least 5 experiments, which we set as a
274 minimum threshold. *Chaetoceros* responded positively to high CO₂ in 6 out of 9
275 experiments (Σ RDR of winning = 84; Σ RDR of losing = 61; Fig 5A). *Cylindrotheca*
276 responded positively in 2 out 5 experiments (Σ RDR of winning = 5; Σ RDR of losing =
277 9; Fig. 5B). *Pseudo-Nitzschia* responded positively in 2 out of 9 experiments (Σ RDR of
278 winning = 3; Σ RDR of losing = 77; Fig. 5C). Thus, *Pseudo-Nitzschia* is the only genus,
279 for which there seems to be a fairly consistent negative response to high CO₂.

280 4. Discussion

281 Numerous physiological studies have shown that diatom growth and metabolic rates can
282 be affected by seawater CO₂ concentrations, and that these responses vary widely among
283 different species (Gao and Campbell, 2014). Such inter-specific differences in pCO₂
284 sensitivity are an important feature as this could alter the composition of diatom
285 assemblages in a changing ocean. In this regard, it is interesting to note that
286 paleolimnologists have long been using diatom species composition as paleo-proxy to
287 reconstruct lake pH (Battarbee et al., 2010). Hence, there is ample evidence that high CO₂
288 conditions have the potential to change the diatom species composition.



289 Indeed, our analysis revealed that CO₂-induced changes in diatom community
290 composition occurred in 27 out of 40 (i.e. 68 %) of community-level experiments which
291 investigated species composition (Fig. 4C). This is certainly a conservative outcome
292 because many studies have only looked at dominant species. In fact, one of the few
293 experiments that investigated the diatom assemblage with higher taxonomical resolution
294 found CO₂ effects also on subdominant species (Sommer et al., 2015) which may have
295 been overlooked in many other experiments.

296 **4.1 Winners and losers in the diatom community**

297 There was sufficient data (i.e. ≥ 5 experiments) for the genera *Chaetoceros*,
298 *Cylindrotheca*, and *Pseudo-Nitzschia* to determine common responses to high CO₂.
299 Among these 3, only *Pseudo-Nitzschia* was fairly consistently identified as a “loser”
300 within the investigated natural diatom communities. *Chaetoceros* was mostly winning
301 while *Cylindrotheca* was mostly losing but the trends were not strong. The relatively
302 weak performance of *Pseudo-Nitzschia* spp. was somewhat surprising because previous
303 monoclonal experiments with this genus often reported a sometimes pronounced positive
304 (Sun et al., 2011; Tatters et al., 2012), or no influence of high CO₂ on their growth rate
305 (Sugie and Yoshimura, 2013; Trimborn et al., 2013) but more rarely a negative one
306 (Tatters et al., 2013). Likewise, laboratory competition experiments between *Chaetoceros*
307 *debilis* and *Pseudo-Nitzschia subcurvata* saw the latter rather on the winning side under
308 high CO₂ although the difference between them was small (Trimborn et al., 2013). The
309 reasons for the inconsistency between our results and the impression derived from
310 controlled laboratory experiments could be manifold. Since our outcome is based on
311 “only” 9 experiments, it could still be coincidence. However, the pronounced difference
312 in the RDR value alleviate this concern to some extent (see numbers on top of Fig. 5C).
313 If the inconsistency is reflecting a true biological pattern than this would emphasize once



314 more that ecological success within a natural community cannot be easily derived from
315 physiological studies.

316 **4.2 CO₂ effects on diatom assemblages originating from (direct) physiological**
317 **responses to high CO₂**

318 Most studies that found effects of pCO₂ on diatom communities related these changes to
319 CO₂ fertilization of photosynthesis. Concentrations of CO₂ in the surface ocean are
320 relatively low compared to other forms of inorganic carbon, especially bicarbonate ion
321 (HCO₃⁻) (Zeebe and Wolf-Gladrow, 2001). However, RubisCO, the primary
322 carboxylating enzyme used in photosynthesis, is restricted to CO₂ for carbon fixation and
323 has a relatively low affinity for CO₂ compared to O₂ (Falkowski and Raven, 2007).
324 Therefore, diatoms (like many other phytoplankton species) operate a carbon
325 concentrating mechanism (CCM) to enhance their CO₂ concentration at the site of
326 fixation relative to external concentrations (e.g. by converting HCO₃⁻ to CO₂) and thereby
327 establish higher rates of carbon fixation than what would be possible when only
328 depending on diffusive CO₂ uptake (Giordano et al., 2005). It is well known that the
329 proportion of CO₂ uptake vs. HCO₃⁻ uptake for photosynthesis varies largely among
330 diatoms (Burkhardt et al., 2001; Rost et al., 2003; Trimborn et al., 2008) and is
331 theoretically also a function of cell size (Flynn et al., 2012; Wolf-Gladrow and Riebesell,
332 1997). Accordingly, increasing seawater pCO₂ may increase the proportion of diffusive
333 carbon uptake and/or lower the energy and resource requirements for CCM operation
334 (Raven et al., 2011). From a physiological point of view, these mechanisms could allow
335 for increased rates of photosynthesis and cell division.

336 So how do these theoretical considerations align with (A) the variable and species-
337 specific physiological responses of diatoms to increasing CO₂ (Dutkiewicz et al., 2015),



338 and (B) the results from community-level experiments compiled in this study? Regarding
339 the variability of physiological responses, progress has recently been made by (Wu et al.,
340 2014) who experimentally demonstrated a positive relationship between cell volume and
341 the magnitude of the CO₂ fertilization effect on diatom growth rates. Their findings agree
342 well with theoretical considerations, which predict that high CO₂ is particularly beneficial
343 for carbon acquisition by larger species as they are more restricted by diffusion gradients
344 due to lower surface-to-volume ratios than smaller cells (Flynn et al., 2012; Wolf-
345 Gladrow and Riebesell, 1997). The outcome of our literature analysis supports this
346 allometric concept (Fig. 4, Table 2). Twelve out of 13 experiments in which cell size was
347 taken into account found a shift towards larger species. This is reflected in the Σ RDR
348 score of 265 which is ~25 times higher than the opposite result (i.e. CO₂-induced shifts
349 towards smaller diatoms, Fig. 4C). An allometric scaling of CO₂ sensitivity is particularly
350 useful for modelling since cell size is a universal trait which is relatively easy to measure
351 and therefore frequently available (Ward et al., 2012). Accordingly, it may lead to
352 significant improvements of ecological and/or biogeochemical model projections under
353 CO₂ forcing when more than one size class for diatoms is considered.

354 However, although the (Wu et al., 2014) allometric approach constitutes a solid starting
355 point to help understanding the variable responses of different diatom species, it probably
356 also still needs some further refinements. For example, central components of CCMs
357 seem to be adapted to diatom cell sizes, thereby potentially alleviating a strict cell size
358 dependency of CO₂ limitation (Shen and Hopkinson, 2015). Furthermore, size
359 dependency alone cannot account for taxon-specific differences in the mode of carbon
360 acquisition (diffusive uptake of CO₂ vs. CCM-supported uptake of HCO₃⁻) and how this
361 will affect the competitive ability of species under increasing CO₂. OA will lead to much
362 larger changes in dissolved CO₂ than in HCO₃⁻. Thus, species that rely to a larger extent



363 on a resource-intensive CCM may benefit more from increasing pCO₂ on a cellular level,
364 as they could increase the proportion of diffusive CO₂ uptake. However, it is also possible
365 that the same species would be disadvantaged on the community-level, because their
366 niche (that is, being competitive at lower CO₂ due to an efficient CCM) is diminished
367 under high CO₂ conditions. Which of the scenarios occurs in nature would also depend
368 on how flexible species are in terms of switching carbon acquisition modes, as well as
369 resource allocation. In this regard, it is noteworthy that only few physiological studies on
370 OA effects have taken into account the role of changing nutrient concentrations or even
371 a transition to nutrient limitation. The available experimental evidence suggests that
372 increasing pCO₂ may reduce cellular nutrient requirements for CCM operations and
373 therefore free resources for elevated maximum diatom population densities, particularly
374 when running into nutrient limitation (Taucher et al., 2015). Unfortunately, however, the
375 relevance of this mechanism has so far only been investigated in monoclonal laboratory
376 experiments but not on the community-level.

377 These considerations illustrate that cell size is an important factor, but is not sufficient to
378 predict physiological or even community-level of diatoms to OA. Moreover, the
379 allometric concept as well as the additional mechanisms described above generally
380 presume positive effects of CO₂-fertilization, thus yielding no first order explanations for
381 observed negative responses of diatoms to changing carbonate chemistry. Obviously,
382 increasing CO₂ concentrations are accompanied by increasing proton (H⁺) concentrations
383 under ocean acidification. High H⁺ concentrations may reduce key metabolic rates above
384 certain thresholds and outweigh the positive influence of CO₂ fertilization as has been
385 observed in coccolithophores (Bach et al., 2011, 2015).

386 Another pathway by which ocean acidification may alter diatom communities is the pH
387 effect on silicification and silica dissolution. Low seawater pH should theoretically



388 facilitate silicification as the precipitation of opal occurs in a cellular compartment with
389 low pH conditions (pH ~5) (Martin-Jézéquel et al., 2000; Vrieling et al., 1999). At the
390 same time, a lower pH should reduce chemical dissolution rates of the SiO₂ frustule
391 (Loucaides et al., 2012). While experimental evidence on this topic is still scarce and
392 partly controversial (Hervé et al., 2012; Mejía et al., 2013; Milligan et al., 2004), it is not
393 unlikely that OA-induced changes in the formation and dissolution of biogenic silica may
394 alter the strength of the frustule and therefore the palatability of diatoms to zooplankton
395 grazers (Friedrichs et al., 2013; Hamm et al., 2003; Liu et al., 2016; Wilken et al., 2011).
396 As for the other physiological effects e.g. on carbon fixation, it is likely that OA impacts
397 on silicification will vary among different diatoms species e.g. according to their species-
398 specific intrinsic buffering capacity, thereby leading to further taxonomic shifts within
399 diatom communities.

400 The response of diatoms to increasing pCO₂ in natural environments will be further
401 modified by multiple other environmental drivers changing simultaneously. Climate
402 change is expected to elevate ocean temperature, as well as also irradiance and nutrient
403 availability via changes in stratification. Physiological experiments have shown that
404 elevated pCO₂ may have beneficial effects under low and moderate irradiance, but this
405 effect may reverse under high light conditions due to enhanced photoinhibition (Gao
406 2012). Analogously, warming may have positive or negative effects on photosynthesis
407 and metabolism in general, depending on the thermal optima of the respective species
408 (Boyd et al., 2018). Altogether, these multiple additional drivers will also affect diatom
409 communities, leading to shifts in their taxonomic composition and size structure, which
410 will interact with the impacts of OA.

411 **4.3 Indirect CO₂ effects on diatom assemblages through food web interactions**



412 Diatom community responses can not only originate from a direct CO₂ effect on their
413 physiology but also be caused indirectly through CO₂ responses on other components of
414 the food web. For example, if a grazer of a diatom species is negatively affected by OA
415 then this may benefit the prey and indirectly promote its abundance. Direct OA impacts
416 on zooplankton communities are usually assumed to play a minor role, although there is
417 some experimental evidence that lower pH may have physiological effects at least on
418 some sensitive species or developmental stages (Cripps et al., 2016; Thor and Dupont,
419 2015; Thor and Oliva, 2015). Nevertheless, much of the currently available empirical
420 evidence indicates that zooplankton communities are affected by OA rather via bottom-
421 up effects, e.g. via changes in primary production or taxonomic composition of the
422 phytoplankton community (Meunier et al., 2017). However bottom-up effects on
423 zooplankton biomass, size structure, or species composition may in turn trigger feedbacks
424 on diatom communities, thereby leading to a feedback loop that may reinforce until a new
425 steady state is reached. Such considerations illustrate that also second or third order
426 effects need to be considered when assessing OA effects on the level of ecological
427 communities. Accounting for such indirect effects requires a holistic approach
428 considering all key players in of the food web (something that is beyond the scope of this
429 study). Therefore, interpretations about what the observed responses could mean for
430 entire plankton food webs or even biogeochemical element cycles (section 4.4) should
431 always be regarded with some healthy skepticism as they often neglect the potential for
432 indirect effects.

433 **4.4 Implications of changes in diatom community structure for pelagic food webs** 434 **and biogeochemical cycles**

435 The taxonomic composition and size structure of phytoplankton communities influences
436 the transfer of energy from primary production to higher trophic levels. In theory, larger



437 diatoms should support a more direct transfer because less trophic intermediates are
438 needed and therefore less respiration occurs until prey items are in an appropriate size
439 range for top predators (Azam et al., 1983; Pomeroy, 1974; Sommer et al., 2002).
440 Likewise a reduced abundance of the potentially toxic genus *Pseudo-Nitzschia* under high
441 CO₂ could further improve trophic transfer and growth of consumers in food webs where
442 *Pseudo-Nitzschia* exerts harmful impacts at present. Such changes at the bottom of a food
443 web might eventually lead to higher production in higher trophic levels such as fish.
444 Indeed, recent experimental evidence indicated that fish (including commercially
445 important species) could under certain constellations benefit from high CO₂ due to higher
446 food availability, although it was not tested if this response is somehow linked to the
447 diatom community (Goldenberg et al., 2018; Sswat et al., 2018).

448 Fluxes of elements through the oceans are (like fluxes of energy through food webs)
449 influenced by the composition of diatom communities (Sarmiento and Gruber, 2006).
450 This is particularly well recognized in the context of organic carbon export to the deep
451 ocean, for which diatoms are considered to play a pivotal role (Smetacek, 1985). Given
452 that high CO₂ favours large and perhaps more silicified diatoms over smaller ones
453 (section 4.2), we might expect accelerated sinking and thus a positive feedback on the
454 vertical carbon flux. This classical hypothesis is supported by observational evidence
455 from two consecutive years of the North Atlantic spring bloom where, despite similar
456 primary production, export was much higher in the year when the larger diatom species
457 dominated (Boyd and Newton, 1995). However, whether the positive relationship
458 between size and carbon export holds under all circumstances is by no means clear
459 (Tréguer et al., 2018). It is possible that shifts towards larger sized species coincide with
460 shifts in other traits that feed back negatively on carbon export. For example, when the
461 size shift is associated with decreasing C:Si stoichiometry it may ultimately reduce carbon



462 export (Assmy et al., 2013).

463 The abovementioned examples of trophic transfer and export fluxes illustrate the
464 importance of the factor “diatom community structure” in the context of marine food
465 production and biogeochemical fluxes. They also illustrate that our understanding of the
466 feedbacks induced through changes in diatom communities is highly incomplete. Hence,
467 with our limited understanding we must currently classify CO₂-induced changes in
468 diatom communities as “a potential risk” causing changes in key ecosystem services.

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948 **Tables and Figures**

949

950 **Table 1.** Response of diatom communities to high CO₂. 69 experiments from 54
951 publications were considered here. Location refers to the place where diatom
952 communities were collected. The RDR is dimensionless (see methods). T is the average
953 incubation temperature in °C. DoE are days of experiment with the number of samplings
954 given as the second number. Pre-filt. gives the mesh size in case the collected plankton
955 community was pre-filtered before incubation. Setup refers to the incubation style:
956 undiluted volumes (batch), repeatedly diluted volumes (s.-cont.), flow-through setups
957 (fl.-thr.; only benthos), chemostats (chem.; only pelagic), CO₂ vent sites (seep; only
958 benthos). Incubations can either be performed on deck (e.g. shipboards), *in situ* (e.g. *in*
959 *situ* mesocosms) or under laboratory conditions. V refers to the incubation volume.
960 Nutrient amendments were made in some but not all studies. The element indicates
961 which nutrients were added. Asterisks indicate the presense of residual nutrients at the
962 beginning of the study. Manipulations were done with: CO₂ saturated seawater (SWsat),
963 acid additions (Acid), combined additions of acid and base (Comb.), CO₂ gas additions
964 (CO₂), Aeration at target *p*CO₂ (Aer.), Passing CO₂ gas through a diffusive silicone tubing
965 (Diff.). Meth. indicates the applied methodology to investigate diatom communities: light
966 microscopy (LM), pigment analyses (PA), flow cytometry (FC), genetic tools (PCR),
967 biogenic silica (BSi). The *p*CO₂ range of the experiment with the number of treatments
968 given in brackets. The response of the bulk diatom community to CO₂: no effect (~),



969 positive (p), negative (n), not reported (N/A). The pCO₂ response indicates approximately
 970 in between which treatments a CO₂ response was observed. Please note that this is based
 971 on visual inspection of the datasets and therefore involves subjectivity. Please also note
 972 that the range equals the treatment values in case only two treatments were set up. CO₂
 973 induced shifts between diatom species can be: shift to larger species (large), shift to
 974 smaller species (small), unspecified shift (shift), no species shift detected (~), not reported
 975 (N/A). Winners or losers of the diatom community comprise: *Chaetoceros* (Chae), large
 976 *Chaetoceros* (Chae I), medium *Chaetoceros* (Chae II), small *Chaetoceros* (Chae III),
 977 *Neosynadra* (Neos), *Rhabdonema* (Rhab), *Eucampia* (Euca), *Cerataulina* (Cera),
 978 *Thalassiosira* (Thals), *Proboscia* (Prob), *Pseudo-nitzschia* (Ps-n), *Thalassionema*
 979 (Thalns), *Cylindrotheca* (Cyli), *Guinardia* (Guin), *Synedropsis* (Syned), *Dactyliosolen*
 980 (Dact), *Toxarium* (Toxa), *Leptocylindrus* (Lept), *Grammatophora* (Gram), *Bacillaria*
 981 (Baci), *Navicula* (Navi).

Reference	lat	long	RDR	S	T (°C)	Habitat	DoE/ # of sampl.	Pre-filt. (µm)	Setup	Incub.	V (L)	Nutr.	Manip.	Meth.	pCO ₂ range (µatm)	CO ₂ effect	pCO ₂ response (µatm)	Intra-taxon effect	Winners	Losers
(Bach et al., 2017)	58.264	11.479	76.2	29	7	est.	113/57	1000	batch	in situ	50000	*none	SWsat	PA, LM	(2) 380, 760	p	380 - 760	large	Cosc	
(Bach et al., 2019)	27.990	15.369	59.6	37	18.5	coastal	32/21	3000	batch	in situ	8000	N,P,Si	SWsat	LM, BSi	(7) 380 - 1120	p	380 - 1120	large	Chae, Guin, Lept	Nitz
(Biswas et al., 2011)	16.750	81.100	2.1	25	29.5	est.	5/2	200	batch	Deck	5.6	*none/N, P	Comb.	PA	(4) 230 - 1860	n	650 - 1400	N/A		
(Biswas et al., 2017)	17.000	83.000	1.5	?	?	coastal	2/1	200	batch	Deck	2	*N,P,Si,F e,(Zn)	Comb.	LM	(2) 230, 2200	p	230 - 2200	shift	Skel	Thals
(Davidson et al., 2016)	-68.583	77.967	10.5	34	0.1	coastal	8/5	200	batch	Lab	650	*Fe	SWsat	LM	(6) 80 - 2420	n	1280 - 1850	small	Frag	Chae
(Domingues et al., 2017)	37.017	-8.500	7.4	?	23.5	est.	1/1	no	batch	Deck	4.5	N,P,Si,NH 4	Comb.	LM, PA	(2) 420, 710	~		~		
(Donahue et al., 2019)	-45.800	171.130	2.6	34	11	oceanic	14/5	200	batch	Lab	10	*Fe	Diff.	LM, FC	(2) 350, 620	~		N/A		
(Donahue et al., 2019)	-45.830	171.540	2.6	34	11	oceanic	21/4	200	batch	Lab	10	*Fe	Diff.	LM, FC	(2) 350, 630	p		N/A		
(Eggers et al., 2014)	38.633	27.067	1.9	36	15	coastal	9-10/3	200	batch	Deck	4	N,P,Si	Comb.	LM	(2) 380, 910	p	380 - 910	large	Chae III	Thals
(Eggers et al., 2014)	38.650	27.250	1.9	36	15	coastal	9-10/4	200	batch	Deck	4	N,P,Si	Comb.	LM	(2) 380, 910	p	380 - 910	large	Thals, Chae II	Chae I
(Eggers et al., 2014)	38.617	27.250	1.9	36	15	oceanic	9-10/5	200	batch	Deck	4	N,P,Si	Comb.	LM	(2) 380, 910	~		N/A		



(Endo et al., 2013)	46.000	160.000	2.8	33	14	oceanic	14/3	197	batch	Deck	12	*none	Aer.	PA	(4) 230 - 1120	~		N/A		
(Endo et al., 2015)	53.083	177.000	2.8	?	8.2	oceanic	5/3	197	batch	Deck	12	*none	Aer.	PA, PCR	(2) 360, 600	n	360 - 600	~		
(Endo et al., 2016)	41.500	144.000	2.8	?	5.4	oceanic	3/3	197	batch	Deck	12	*Fe	Aer.	PA, PCR	(4) 180 - 1000	n	350 - 1000	shift		
(Feng et al., 2009)	57.580	15.320	1.7	35	12	oceanic	14/1-2	200	s.-cont.	Deck	2.7	N,P	Aer.	LM, PA	(2) 390, 690	p	390 - 690	large	Ps-n	Cyli
(Feng et al., 2010)	-74.230	179.230	1.7	34	0	oceanic	18/1-14	200	s.-cont.	Deck	2.7	none	Aer.	LM, PA	(2) 380, 750	~		large	Chae	Cyli
(Gazeau et al., 2017)	43.697	7.312	125.8	38	14	coastal	18/14	5000	batch	in situ	45000	none	SWsat	PA	(6) 350 - 1250	p	600 - 1000	N/A		
(Gazeau et al., 2017)	42.580	8.726	125.8	38	23	coastal	27/18	5000	batch	in situ	45000	none	SWsat	PA	(6) 420 - 1250	~		N/A		
(Grear et al., 2017)	41.575	71.405	9.3	?	9	est.	6/7	no	chem.	Deck	9.1	?none	Comb.	LM	(3) 220 - 720	~		~		
(Hama et al., 2016)	34.665	138.940	7.1	?	?	coastal	29/11	100	batch	Deck	400	N,P,Si	Aer.	PA	(3) 400 - 1200	~		N/A		
(Hare et al., 2007)	56.515	164.730	6.0	?	10.4	coastal	9-10/5	no	s.-cont.	Deck	2.5	Fe,N,P,Si	Aer.	LM, PA	(2) 370, 750	n	370 - 750	shift		Cyli
(Hare et al., 2007)	55.022	179.030	6.0	?	10.4	oceanic	9-10/3	no	s.-cont.	Deck	2.5	Fe	Aer.	LM, PA	(2) 370, 750	n	370 - 750	N/A		
(Hopkins et al., 2010)	60.300	5.200	99.1	?	10	coastal	21/9	no	batch	in situ	11000	N, P	Aer.	LM	(2) 300, 600	n	300 - 600	N/A		
(Hoppe et al., 2013)	-66.833	0.000	1.9	34	3	oceanic	27-30/1	200	s.-cont.	Lab	4	*none	Aer.	LM	(3) 200 - 810	N/A	400 - 810	shift	Syned	Ps-n
(Hoppe et al., 2017b)	71.406	68.601	1.9	33	9.5	oceanic	8/3	100	s.-cont.	Deck	8	N,P,Si	Aer.	PA, LM	(2) 320, 990	~		~		
(Hoppe et al., 2017a)	63.964	60.125	1.9	32	7.9	oceanic	13-14/3	100	s.-cont.	Deck	8	N,P,Si	Aer.	LM	(2) 300, 960	n	300 - 960	shift	Frag	Ps-n
(Husserr et al., 2017)	71.406	70.188	2.6	33	4.3	oceanic	9/3-9	200	batch	Deck	10	*none	Comb.	LM, PA	(6) 510 - 3300	n	1040 - 1620	~		
(James et al., 2014)	-45.639	170.671	?	?	11.6	benthic	42/2	fl.-thr.	Lab		0	none	Comb.	pic	(2) 400, 1250	~		N/A		
(Johnson et al., 2011)	38.417	14.950		38	23.5	benthic	21/1	NA	seep	in situ	0	none	NA	PA, LM	(3) 420 - 1600	p	420 - 590	large	Toxa, Gram, Bac, Navi, Cocc	Cycl, Neos, Rhab, Nitz
(Kim et al., 2006)	34.600	128.500	4.3	?	14	coastal	14/7	60	batch	in situ	150	N,P	Aer.	LM	(3) 250 - 750	N/A	400 - 750	shift	Skel	Nitz
(Kim et al., 2010)	34.600	128.500	52.1	?	12	coastal	20/22	no	batch	in situ	1600	N,P,Si	SWsat/Aer.	LM	(2) 400, 900	~		shift	Skel	Euca
(Mallozzi et al., 2019)	29.241	90.935	2.4	12	21	est.	112/9	80	s.-cont.	Lab	20	*none	Aer.	PA, LM	(2) 400, 1000	~		shift	Cyli	
(Mallozzi et al., 2019)	29.272	89.963	2.4	17	21	est.	112/9	80	s.-cont.	Lab	20	*none	Aer.	PA, LM	(2) 400, 1000	~		shift	Cyli	
(Maugendre et al., 2015)	43.667	-7.300	1.9	?	15	oceanic	12/4	200	batch	Deck	4	none	SWsat	PA	(2) 360, 630	~		N/A		
(Nielsen et al., 2010)	56.057	12.648	1.6	19	10.7	est.	14/4	175	s.-cont.	Lab	2.5	*none	Acid	LM, PA	(3) 500 - 1500	~		~		
(Nielsen et al., 2012)	-42.887	147.339	1.8	31	16	coastal	14/4	250	s.-cont.	Lab	2.5	*none	Acid	LM, PA	(3) 300 - 1200	~		~		
(Park et al., 2014)	34.600	128.500	59.6	?	17	coastal	19/17	no	batch	in situ	2400	N,P,Si	SWsat/Aer.	LM, PA	(6) 160 - 830	p	160 - 830	N/A	Cera	



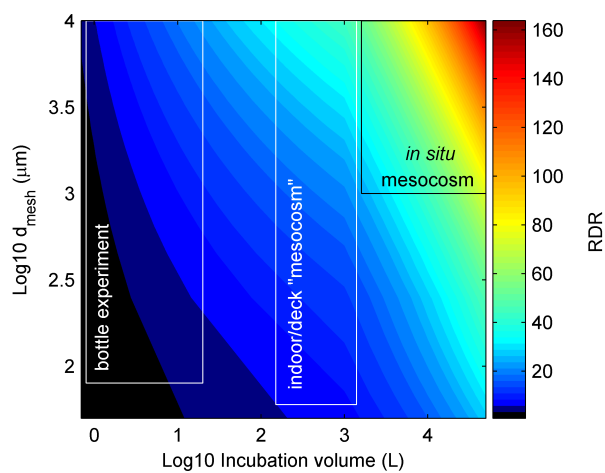
(Paul et al., 2015)	59.858	23.258	112.7	6	11	est.	46/22	3000	batch	in situ	54000	none	SWsat	PA	(6) 370 - 1230	p	820 - 1000	N/A		
(Reul et al., 2014)	36.540	-4.600	3.3	?	21	coastal	7/6	200	batch	Deck	20	control/N _P	Aer.	LM, PA	(2) 500, 1000	p	500 - 1000	large		
(Roleda et al., 2015)	-45.639	170.671		34	10.8	benthic	112/7	NA	fl-thr.	Lab	0.65	none	Comb.	PA	(2) 430, 1170	~		N/A		
(Rossoll et al., 2013)	54.329	10.149	29.8	18	18	est.	28/7	no	batch	Lab	300	N,P,Si	Aer.	LM	(5) 390 - 4000	~		N/A		
(Sala et al., 2015)	41.667	2.800	26.1	38	14	coastal	9/2	no	batch	Lab	200	none	CO2	LM	(2) 400, 800	~		N/A		
(Sala et al., 2015)	41.667	2.800	26.1	38	22	coastal	9/2	no	batch	Lab	200	none	CO2	LM	(2) 400, 800	~		N/A		
(Schulz et al., 2008)	60.267	5.217	133.7	31	10.5	coastal	25/18-23	no	batch	in situ	27000	N,P	Aer.	PA	(3) 350 - 1050	~		N/A		
(Schulz et al., 2013)	78.937	11.893	158.5	34	3	coastal	30/26 - 30	3000	batch	in situ	45000	N,P,Si	SWsat	LM, PA	(8) 185 - 1420	~		N/A		
(Schulz et al., 2017)	60.265	5.205	125.8	32	9	coastal	38/35	3000	batch	in situ	75000	*N,P	SWsat	LM, PA	(8) 310 - 3050	n	1165 - 1425	N/A		
(Segovia et al., 2017)	60.390	5.320	99.1	?	11	coastal	22/9	no	batch	in situ	11000	control	SWsat /Aer.	FC	(2) 300, 800	~		N/A		
(Sett et al., 2018)	54.329	10.149	49.8	20	5	est.	44/26	sand filter	batch	Lab	1400	*none	SWsat	LM, FC	(2) 540, 1020	~		~		
(Shaik et al., 2017)	15.453	43.801	5.6	35	29	coastal	2/1	no	batch	Deck	2	N,P,Si,Fe	CO2	LM	(2) 330, 1000	p	330 - 1000	~		
(Shaik et al., 2017)	15.453	43.801	5.6	36	29	coastal	9/1	no	s-cont.	Deck	2	N,P,Si,Fe	CO2	LM	(2) 400, 1000	p	400 - 1000	~		
(Shaik et al., 2017)	15.453	43.801	5.6	35	29	coastal	2/1	no	batch	Deck	2	N,P,Si,Fe	CO2	LM	(2) 240, 780	p	240 - 780	~		
(Sommer et al., 2015)	54.329	10.149	49.8	20	9.15	est.	24/11	sand filter	batch	Lab	1400	*none	SWsat	LM	(2) 440, 1040	~		shift		Prob, Thal, Guin, Ps-n, Chae
(Tatters et al., 2013)	-45.752	170.810	0.8	35	14	coastal	14/2	80	s-cont.	Lab	0.8	N,P,Si,Fe	Aer.	LM	(3) 230 - 570	N/A	400 - 570	shift	Cosc, Ps-n	Navi, Chae
(Tatters et al., 2018)	33.750	118.215	12.1	?	19	coastal	10/1	no	chem.	Deck	20	N/urea,P, Si	Aer.	LM	380, 800	N/A		shift		
(Taugher et al., 2018)	27.928	15.365	97.6	37	24-22	coastal	60/35	3000	batch	in situ	35000	N,P,Si	SWsat	LM, PA	(8) 350 - 1030	p	890 - 1030	large	Guin	Lept
(Thoisen et al., 2015)	69.217	53.367	1.4	33	3	coastal	8-17/6-9	250	s-cont.	Lab	1.2	*none	SWsat	LM	(4) 440 - 3500	n	440 - 900	shift	Navi I	Navi II
(Tortell et al., 2002)	-6.600	81.017	7.1	?	?	oceanic	11/4	no	s-cont.	Deck	4	*none	Aer.	PA, LM	(2) 150, 750	p	150 - 440	~		
(Tortell et al., 2008)	NA	NA	7.1	?	0	N/A	10-18/7	no	s-cont.	Lab	4	*Fe	Aer.	LM, PA	(3) 100 - 800	p	100 - 400	large	Chae	Ps-n
(Tortell et al., 2008)	NA	NA	7.1	?	0	N/A	10-18/7	no	s-cont.	Deck	4	*Fe	Aer.	LM, PA	(3) 100 - 800	p	100 - 400	large	Chae	Ps-n
(Tortell et al., 2008)	NA	NA	7.1	?	0	N/A	10-18/7	no	s-cont.	Deck	4	*Fe	Aer.	LM, PA	(3) 100 - 800	p	100 - 400	large	Chae	Ps-n
(Trimborn et al., 2017)	-53.013	10.025	1.9	34	3	oceanic	30/4	200	s-cont.	Lab	4	none	Aer.	LM	420, 910	n	420 - 910	shift		Ps-n
(Witt et al., 2011)	-23.450	151.917		?	24-25	benthic	11/4	NA	fl-thr.	Deck	10	none	SWsat	LM	(4) 310 - 1140	p	560 - 1140	N/A		
(Wolf et al., 2018)	78.917	11.933	1.9	?	3	coastal	10 - 13/1	200	s-cont.	Lab	4	none	Aer.	LM	(2) 400, 1000	N/A		~		
(Yoshimura et al., 2010)	49.500	148.250	2.7	33	13.5	oceanic	14/5	243	batch	Deck	9		Aer.	PA	(4) 150 - 590	n	150 - 280	N/A		



(Yoshimura et al., 2013)	53.390	177.01 0	2.8	?	8.4	oceanic	14/3	197	batch	Deck	12	*none	Aer.	PA, LM	4 (300 - 1190)	p	960 - 1190	N/A		
(Yoshimura et al., 2013)	49.020	174.02 0	2.8	?	9.2	oceanic	14/3	197	batch	Deck	12	*none	Aer.	PA, LM	(4) 230 - 1110	p	880 - 1110	N/A		
(Young et al., 2015)	-44.779	64.073	7.1	?	-1	coastal	21/21	no	s- cont.	Deck	4	*none	Aer.	PA	(3) 100 - 800	~		N/A		
(Young et al., 2015)	-44.780	64.073	7.1	?	-0.5	coastal	16/16	no	s- cont.	Deck	4	*none	Aer.	PA, LM	(3) 100 - 800	~		N/A		
(Young et al., 2015)	-44.780	64.073	7.1	?	1.5	coastal	20/20	no	s- cont.	Deck	4	*none	Aer.	PA	(3) 100 - 800	~		N/A		

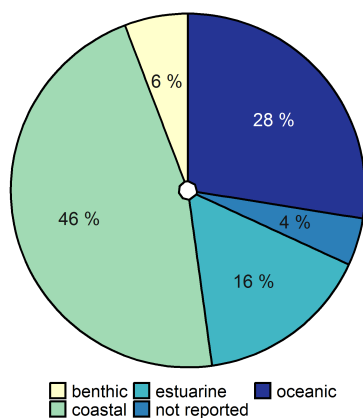
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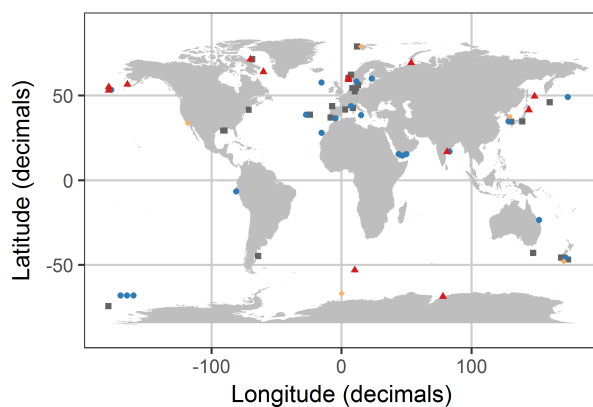
984

985 **Figure 1.** RDR as a function of incubation volume and size of the mesh that was used
 986 while filling the incubation volumes (d_{mesh}). The black and white boxes illustrate
 987 approximate ranges of the three main types of containers used in experiments. Please note
 988 that the general definition for mesocosms are volumes >1000 L (Guangao, 1990) but
 989 since most authors also use this term for open batch incubations with volumes between
 990 150 – 1000 L we also stick to this term for the intermediate class.



991

992 **Figure 2.** Habitats in which the ocean acidification experiments were conducted. The
993 total number of studies is 69. ‘not reported’ means that coordinates where the incubation
994 water was collected were not provided.

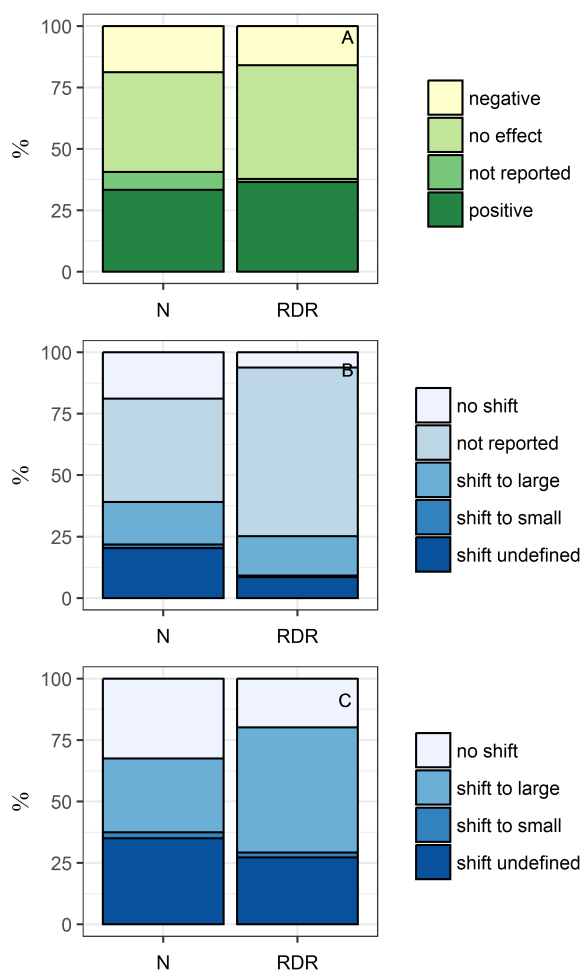


995

996 **Figure 3.** Distribution of diatom experiments with associated OA response of the bulk
997 communities as listed in Table 1. Blue circles = positive effect; red triangles = negative
998 response; grey squares = no response; orange diamonds = response not reported.
999 Locations were slightly modified in case of geospatial overlap to ensure visibility. Please



1000 note that the three blue points in the Ross Sea at about -68, -165 are approximate locations
 1001 because the reference did not provide coordinates.

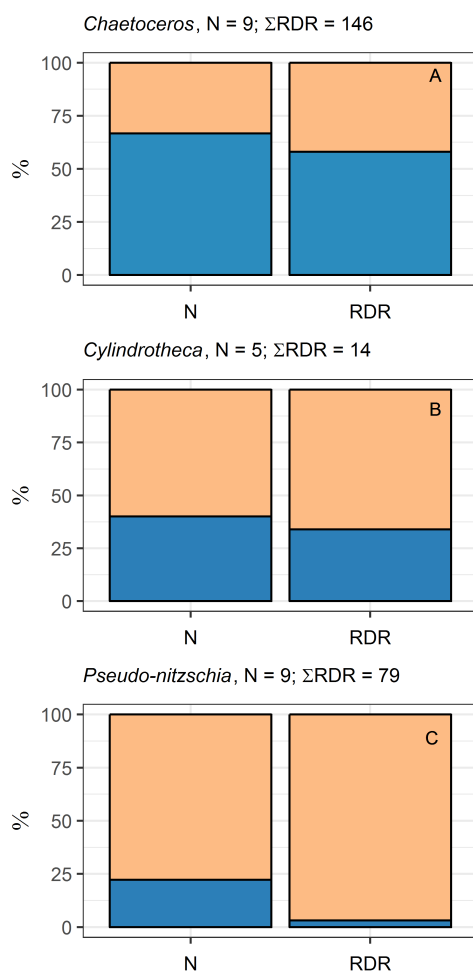


1002
 1003 **Figure 4.** Summary of the literature analysis. (A) Response of the bulk diatom
 1004 community to ocean acidification. (B) Shifts among different diatom species due to ocean
 1005 acidification. ‘Shift to large’ and ‘shift to small’ indicate that the diatom community
 1006 shifted towards the dominance of larger or smaller species, respectively. (C) Same data
 1007 as in B but excluding studies where species shifts within the diatom community were not
 1008 reported. This reduced the dataset from 69 to 40 studies. The left column is based on the



1009 number of studies. For example, the bulk diatom community was positively affected by
1010 OA in 29 out of 69 studies which is 33 %. The right column is based on the RDR values.
1011 For example, the Σ RDR value of all studies where the diatom community was positively
1012 affected by OA was 605 which is 36 % of the total Σ RDR.

1013



1014

1015 **Figure 5.** Winners and losers in diatom communities. The bar chart indicates the fraction
1016 of experiments where the respective genera benefitted from high CO₂ (blue) or were put
1017 at a disadvantage relative to the control treatment (orange). Shown here are diatom



1018 genera that were microscopically identified in at least 5 experiments. The left bars are
1019 fractions based on the number of experiments (N, total number given above each plot).
1020 The right bars are fractions based on the RDR values of these experiments (\sum RDR of all
1021 experiments considered given above each plot). (A) *Chaetoceros*. (B) *Cylindrotheca*. (C)
1022 *Pseudo-nitzschia*. Please note that any such evaluation on the species level cannot be done
1023 at present due to too few data.