



Contents lists available at ScienceDirect

# Renewable and Sustainable Energy Reviews

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## Prediction of diffuse horizontal irradiance using a new climate zone model

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### ARTICLE INFO

#### Keywords:

Diffuse horizontal irradiance  
Global horizontal irradiance  
Direct normal irradiance  
Clearness index  
Diffuse fraction  
Separation method

### ABSTRACT

Knowledge on the diffuse horizontal irradiance (DHI), and direct normal irradiance (DNI) is crucial for the estimation of the irradiance on tilted surfaces, which in turn is critical for photovoltaic (PV) applications and for designing and simulating concentrated solar power (CSP) plants. Since global horizontal irradiance (GHI) is the most commonly measured solar radiation variable, it is advantageous for establishing a suitable method that uses it to compute DHI and DNI. In this way, a new model for predicting the diffuse fraction ( $K_d$ ) based on the climate zone is proposed, using only the clearness index ( $K_t$ ) as the predictor and 1-min resolution GHI data. A review of the literature on models that use hourly and sub-hourly  $K_t$  values to compute  $K_d$  was also carried out, and an extensive performance assessment of both the proposed model and the models from the literature was conducted using ten statistical indicators and a global performance index (GPI). A set of model parameters was determined for each climate zone considered in this study (arid, high albedo, temperate and tropical) using 48 worldwide radiometric stations. It was found that the best overall performing model was the model proposed in this work.

### 1. Introduction

Global horizontal irradiance (GHI) is the most commonly measured solar radiation variable in the ground-based meteorological stations around the world, both in historical datasets and in geographical distribution. Therefore, it is the best dataset available to quantify solar energy resource and assess undergoing or future solar energy projects. On the other hand, information on both diffuse horizontal irradiance (DHI) and direct normal irradiance (DNI) is also crucial to properly design and optimise solar energy systems. In this way, it is advantageous to find a suitable and accurate method based on the GHI measurements to estimate both DHI and DNI, thus enabling the reconstruction of temporal series of these two components in locations where only GHI measurements are available, mainly due to budget limitations and higher requirements for maintenance and calibration procedures. In fact, whereas pyranometer installations are relatively cheap (USD 5–10 K with a data logger), full stations equipped with a sun tracker, pyranometers and a pyrliometer are quite expensive (around USD 30 K) [1]. DHI and DNI data are essential to accurately determine the global solar irradiance on tilted surfaces, for example in sizing and operation of photovoltaic (PV) systems [2]. The models for the diffuse fraction allow to estimate those components based on the GHI and then determine the irradiance on a tilted surface, by opposition to the one-

step methods of converting GHI, as for example the isotropic sky model [3], the Klucher model [4], the Hay-Davies model [5] and the Reindl model [6]. Concentrated Solar Power (CSP) systems mainly use DNI in its energy capturing and conversion processes due to its directional nature and field of view (aperture angle) that depends on the concentration factor [7]. Therefore, the accurate computation of DHI is of vital importance to design, assess the performance and operate such systems [8].

The response of the scientific community for the need of obtaining DHI and DNI data at low cost was given by developing separation models in which the work of Liu and Jordan [9] was the pioneer. That work reported the relation between the clearness index (the ratio between GHI that reaches the surface of the earth and the extraterrestrial irradiance on a horizontal surface,  $K_t$ ) and the diffuse fraction (the ratio of DHI to GHI,  $K_d$ ) using measurements from 98 stations in Canada and United States. The good results obtained by Liu and Jordan lead to the development of several other separation models for different locations. Page [10] developed a model based on monthly mean values for latitudes between 40° N and 40° S. Tuller [11] analysed daily and monthly data to establish models for four locations in Canada. Klein [12] used experimental measurements to assess and validate the model proposed by Liu and Jordan [9] and extended it to allow calculation of monthly average solar irradiation on surfaces with multiple orientations.

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Although several other daily and monthly basis models were presented and are available in the literature, they are not the focus of this work. The work of Khorasanizadeh and Mohammadi [13] reports a comprehensive review of such models. Separation models for high-frequency GHI data are needed until high-resolution DNI measurements are available in a global scale, since required temporal resolution of nowadays reported solar radiation data increased, due to the requirements of high-frequency measurements in the simulation of CSP projects [14]. Therefore, because sub-hourly models are relatively rare in the literature, this work focuses on the available hourly and sub-hourly separation models whose solo predictor is the  $K_t$  and on their ability in representing high-frequency data in a global scale and for different climates. The main reason for using only  $K_t$  as the predictor is due to the greater availability of GHI data worldwide, thus allowing a straightforward evaluation of the model for the higher number of locations as possible. Regarding this type of models that use only  $K_t$  as the predictor, Orgill and Hollands [15] presented a separation model using hourly measurements covering the period from September 1967 to August 1971 for Toronto Airport, Canada. This was the first model found in the literature that met the features mentioned above. In Section 2 are presented all the other models reviewed in this work.

The assessment of new separation models is usually carried out through the comparison of that new model against ground measurements and other models [1] using statistical indicators. Beside some researchers have already presented performance analysis using only models available in the literature [14,16,17], the majority of the validation studies were reported when new models were derived, as is the case of this work. The first hourly models presented [15,18] were compared against the Liu and Jordan monthly model [9]. As time went by, more hourly models became available for test, and therefore models such as the Orgill and Hollands [15] and the Erbs et al. [19] were used in numerous validation studies (e.g. Refs. [6,20,21]). Liu and Jordan's model is still occasionally used with the purpose of presenting a historical comparison of the separation models evolution [22]. Regarding the validation using ground-based measurements, the most used statistical indicators to assess the performance of separation models are the mean bias error (MBE), the root mean square error (RMSE) and the correlation coefficient (R).

One-minute data resolution models are very scarce in the literature. One of the few examples is the work of Engerer [1], which presents a diffuse fraction model based on 1-min clearness index data together with other predictors for southeastern Australia. Gueymard and Ruiz-Arias [14] reported the incapability of hourly models to account for cloud enhancement effects, aiming at the need for reliability in hourly models until more specific minutely models appear in the literature. Therefore, the purpose of this study is to develop a new diffuse fraction model based on 1-min measurements from stations around the globe. Since the model presented by Engerer [1] requires more than one input parameter, the performance assessment of the model developed in this work will be conducted against hourly and sub-hourly models whose only predictor is  $K_t$ . To that end, ten statistical indicators were used, namely the mean bias error (MBE), mean absolute error (MAE), root mean square error (RMSE), mean percentage error (MPE), uncertainty at 95% (U95), relative root mean square error (RRMSE), maximum absolute error (erMAX), correlation coefficient (R) and mean absolute relative error (MARE). These statistical indicators were also combined into a global performance index (GPI). The GPI was used in previous studies in this field by Jamil and Akhtar [23]. Other option to combine different statistical indicators is the combined performance index (CPI), as described by Gueymard [24]. A Taylor diagram and a skill score [25] were also used to provide an additional statistical analysis. In this view, a comprehensive performance assessment of the proposed model as well as of other models in the literature is presented aiming at the identification of the best performing model for the estimation of DHI in a minute resolution all over the world. The organization of this paper is as follows: Section 2 presents a review of the hourly and sub-hourly

models for estimating the diffuse fraction, Section 3 presents the data used in this study and the model development, Section 4 presents the results and discussion, and, finally, conclusions are drawn in Section 5.

## 2. Review of the available models

The models available in the literature were developed using several functional forms, number of predictors and for different time resolutions. The first form to obtain the diffuse fraction was a second degree polynomial as a function of the clearness index, as first proposed by Liu and Jordan [9] in 1960. Later, other models were presented using higher polynomial degrees as well as other functions such as the logistic [26] and the double exponential [27] forms. Several models included other predictors than  $K_t$ , such as sunshine duration, zenith angle, air mass, etc. Regarding time resolution, the available models were proposed to estimate the monthly, daily, hourly and sub-hourly diffuse fraction. In this work, a review of the models that use only  $K_t$  as the predictor with hourly or sub-hourly time resolutions is presented. The authors were able to find 121 different models that met the requirements specified above, although more models may be available in other publications or internal reports and communications that are not readily accessible. In many cases, authors present the same model but for different locations. These models are treated here as unique models when assessing their performance in Section 4. Table 1 presents the models studied in this work. The various locations from which authors used data to develop their models are classified according to the climate region as follows [14]: temperate (TM), arid (AR), tropical (TR) and high albedo (HA).

Based on the information of Table 1, it is notorious that the majority of the models were derived using a polynomial form, followed by the double exponential and logistic functions. This information is also useful to see the distribution of the data used to derive the models according to the climate zone and to identify the length of the datasets, as this is a crucial factor on the determination of the coefficients of the models. Fig. 1 presents the distribution of models as a function of the number of years of the datasets and climate zone. Only the models that presented an explicit dataset length in their respective publications were used to produce Fig. 1. Models based on data from several locations, different climatic zones or models derived from distinct dataset lengths were not considered, hence the 100 models examined in Fig. 1. The higher number of models was developed for the temperate (TM) zone, followed by the arid (AR), tropical (TR) and finally the high albedo (HA) zones. This representativeness of the climate zones is also useful to perceive the distribution of solar radiation measuring stations around the globe and how this may affect the model development. One can conclude that the high albedo and the tropical climate zones are not well represented. However, nowadays there are more meteorological stations in these climate zones (see, e.g. Ref. [69]). The most typical length of the training datasets is two years, followed by one, three and four years. It worths to note that some authors used twenty and even thirty years of data to derive their models (e.g. Refs. [27,61]).

## 3. Model development and test data

### 3.1. Model formulation

The model proposed in this study is based on the correlation of two limiting functions for large and small values of  $K_t$  through the expression disclosed by Churchill and Usagi [70] for the correlation of transfer phenomena, described as follows:

$$Y = (1 + Z^n)^{\frac{1}{n}} \quad (1)$$

where the arbitrary exponent  $n$  needs to be selected in order to correlate those functions accurately [70]. This expression can be used to any phenomenon varying uniformly, for example in heat transfer modelling

**Table 1**  
Review of hourly and sub-hourly  $K_d$  models whose only predictor is  $K_t$ .

Model	Reference	Location	Climate	Data period	Constrains	$K_d$
1	Orgill and Hollands [15]	Toronto, Canada	TM	1967–1971	$K_t \leq 0.35$ $0.35 \leq K_t \leq 0.75$ $K_t > 0.75$	$1.0 - 0.249K_t$ $1.557 - 1.840K_t$ 0.177
2	Bruno [18]	Hamburg, Germany	TM	1973–1974	–	$0.310K_t + 0.139\sin(4.620K_t)$
3	Erbs et al. [19]	Four cities in the United States	Various climates	1974–1976 (Various data periods)	$K_t \leq 0.22$ $0.22 < K_t \leq 0.80$ $K_t > 0.80$	$1.0 - 0.0900K_t$ $0.9511 - 0.1604K_t + 4.3880K_t^2 - 16.6380K_t^3 + 12.3360K_t^4$ 0.165
4	Spencer [28]	Albany, Australia	AR	1973–1977	$K_t < 0.35$ $0.35 \leq K_t \leq 0.75$ $K_t > 0.75$	0.890 $1.414 - 1.736K_t$ 0.110
5	Spencer [28]	Alice Springs, Australia	AR	1974–1977	$K_t < 0.35$ $0.35 \leq K_t \leq 0.75$ $K_t > 0.75$	0.750 $1.183 - 1.444K_t$ 0.110
6	Spencer [28]	Geraldton, Australia	AR	1972–1977	$K_t < 0.35$ $0.35 \leq K_t \leq 0.75$ $K_t > 0.75$	0.850 $1.345 - 1.644K_t$ 0.110
7	Spencer [28]	Guildford, Australia	AR	1975–1977	$K_t < 0.35$ $0.35 \leq K_t \leq 0.75$ $K_t > 0.75$	0.780 $1.254 - 1.595K_t$ 0.060
8	Spencer [28]	Hobart, Australia	TM	1971–1977	$K_t < 0.35$ $0.35 \leq K_t \leq 0.75$ $K_t > 0.75$	0.860 $1.360 - 1.678K_t$ 0.100
9	Spencer [28]	Laverton, Australia	AR	1976–1977	$K_t < 0.35$ $0.35 \leq K_t \leq 0.75$ $K_t > 0.75$	0.860 $1.360 - 1.678K_t$ 0.150
10	Spencer [28]	Melbourne, Australia	AR	1970–1977	$K_t < 0.35$ $0.35 \leq K_t \leq 0.75$ $K_t > 0.75$	0.850 $1.352 - 1.668K_t$ 0.100
11	Spencer [28]	Mildura, Australia	AR	1972–1977	$K_t < 0.35$ $0.35 \leq K_t \leq 0.75$ $K_t > 0.75$	0.870 $1.366 - 1.666K_t$ 0.120
12	Spencer [28]	Mt Gambier, Australia	AR	1974–1977	$K_t < 0.35$ $0.35 \leq K_t \leq 0.75$ $K_t > 0.75$	0.930 $1.450 - 1.744K_t$ 0.140
13	Spencer [28]	Port Hedland, Australia	AR	1974–1977	$K_t < 0.35$ $0.35 \leq K_t \leq 0.75$ $K_t > 0.75$	0.710 $1.142 - 1.431K_t$ 0.070
14	Spencer [28]	Rockhampton, Australia	AR	1974–1977	$K_t < 0.35$ $0.35 \leq K_t \leq 0.75$ $K_t > 0.75$	0.790 $1.245 - 1.527K_t$ 0.100
15	Spencer [28]	Waga Waga, Australia	AR	1974–1977	$K_t < 0.35$ $0.35 \leq K_t \leq 0.75$ $K_t > 0.75$	0.800 $1.280 - 1.605K_t$ 0.080
16	Spencer [28]	Australia (average)	Various climates	1970–1977 (Various data periods)	$K_t < 0.35$ $0.35 \leq K_t \leq 0.75$ $K_t > 0.75$	0.830 $1.321 - 1.624K_t$ 0.100
17	Hawladar [20]	Singapore	TM	1962	$K_t < 0.225$ $0.225 \leq K_t \leq 0.775$ $K_t > 0.775$	0.9150 $1.1389 - 0.9422K_t - 0.3878K_t^2$ 0.2150
18	Ineichen et al. [29]	Geneva, Switzerland	TM	1978–1984	$K_t < 0.15$ $K_t \geq 0.15$	0.98 $0.80 + 2.25K_t - 7.93K_t^2 + 5.26K_t^3$
19	Ineichen et al. [29]	Geneva, Switzerland	TM	1978–1984	$K_t < 0.25$ $0.25 \leq K_t \leq 0.80$ $K_t > 0.80$	1.0 $1.38 - 1.52K_t$ 0.16
20	Ineichen et al. [29]	Geneva, Switzerland	TM	1978–1984	$K_t < 0.25$ $K_t \geq 0.25$	1.0 $1.28K_t - 1.40K_t^2$
21	Muneer et al. [21]	New Delhi, India	TR	1971, 1974	$K_t < 0.175$ $0.175 \leq K_t \leq 0.775$ $K_t > 0.775$	0.9500 $0.9698 + 0.4353K_t - 3.4499K_t^2 + 2.1888K_t^3$ 0.2600
22	Bakhsh et al. [30]	Dharan, Saudi Arabia	AR	1983–1984	$K_t < 0.23$ $0.23 \leq K_t \leq 0.80$ $K_t > 0.80$	$1.0 - 0.220K_t$ $1.235 - 1.260K_t$ 0.225
23	Hollands [31]	Toronto, Canada	TM	1967–1971	–	$[1 - b - \sqrt{(1 - b)^2 - 4ab^2K_t(1 - aK_t)}] / (2abK_t)$ $a = 1.115; b = 0.491$
24	Reindl et al. [6]	Five locations in North America and Europe	Various climates	1979–1982 (Various data periods)	$K_t \leq 0.30$ $0.30 < K_t < 0.78$ $K_t \geq 0.78$	$1.020 - 0.248K_t$ $1.450 - 1.670K_t$ 0.147
25	Al-Rihai [32]	Fudhaliyah, Iraq	AR	1984–1987	$K_t < 0.25$ $0.25 \leq K_t \leq 0.70$ $K_t > 0.70$	0.932 $1.293 - 1.631K_t$ 0.151

(continued on next page)

Table 1 (continued)

Model	Reference	Location	Climate	Data period	Constrains	$K_d$
26	Bourges [33]	37 stations across Europe	TM	At least four years of measurements	$K_t \leq 0.20$ $0.20 < K_t \leq 0.35$ $0.35 < K_t \leq 0.75$ $K_t > 0.75$	1.0 $1.116 - 0.580K_t$ $1.557 - 1.840K_t$ 0.177
27	Chandrasekaran and Kumar [34]	Madras, India	TR	1983–1987	$K_t \leq 0.24$ $0.24 < K_t \leq 0.80$ $K_t > 0.80$	$1.0086 - 0.1780K_t$ $0.9686 + 0.1325K_t + 1.4183K_t^2 - 10.1860K_t^3 + 8.3733K_t^4$ 0.1970
28	Chendo and Maduekwe [35]	Lagos, Nigeria	TM	Two years of measurements	$K_t \leq 0.30$ $0.30 < K_t < 0.80$ $K_t \geq 0.80$	$1.022 - 0.156K_t$ $1.385 - 1.396K_t$ 0.264
29	Maduekwe and Chendo [36]	Lagos, Nigeria	TM	1990–1991	$K_t \leq 0.30$ $0.30 < K_t < 0.80$ $K_t \geq 0.80$	$1.021 - 0.151K_t$ $1.385 - 1.396K_t$ 0.295
30	Lam and Li [37]	Hong Kong, China	TM	1991–1994	$K_t \leq 0.15$ $0.15 < K_t \leq 0.70$ $K_t > 0.70$	0.977 $1.237 - 1.361K_t$ 0.273
31	Hijazin [38]	Amman, Jordan	AR	1985	$K_t < 0.10$ $0.10 \leq K_t \leq 0.80$ $K_t > 0.80$	0.744 $0.842 - 0.977K_t$ 0.060
32	Hijazin [38]	Amman, Jordan	AR	1985	–	$0.847 - 0.985K_t$
33	González and Calbó [39]	Two locations in Iberian Peninsula	TM	1994–1996 (Various data periods)	$0.25 < K_t < 0.75$ $K_t \geq 0.75$	$1.421 - 1.670K_t$ $- 0.043 + 0.290K_t$
34	Boland et al. [26]	Geelong, Australia	TM	67 days	–	$1.0/[1.0 + \exp\{8.645(K_t - 0.613)\}]$
35	Boland et al. [26]	Geelong, Australia	TM	67 days	–	$1.0/[1.0 + \exp\{7.997(K_t - 0.586)\}]$
36	De Miguel et al. [40]	North Mediterranean belt area (11 stations)	TM	1974–1996 (Various data periods)	$K_t \leq 0.21$ $0.21 < K_t \leq 0.76$ $K_t > 0.76$	$0.995 - 0.081K_t$ $0.724 + 2.738K_t - 8.320K_t^2 + 4.967K_t^3$ 0.180
37	Li and Lam [41]	Hong Kong, China	TM	1991–1998	$K_t \leq 0.15$ $0.15 < K_t \leq 0.70$ $K_t > 0.70$	0.976 $0.996 + 0.036K_t - 1.589K_t^2$ 0.230
38	Oliveira et al. [42]	São Paulo, Brazil	TM	1994–1999	$0.17 < K_t < 0.75$	$0.97 + 0.80K_t - 3.00K_t^2 - 3.1K_t^3 + 5.2K_t^4$
39	Ulgen and Hepbasli [43]	Izmir, Turkey	TM	1994–1998	$K_t \leq 0.32$ $0.32 < K_t \leq 0.62$ $K_t > 0.62$	0.6800 $1.0609 - 1.2138K_t$ 0.3000
40	Ulgen and Hepbasli [43]	Izmir, Turkey	TM	1994–1998	$K_t \leq 0.32$ $0.32 < K_t \leq 0.62$ $K_t > 0.62$	0.6800 $0.0743 - 19.3430K_t + 206.9100K_t^2 - 719.7200K_t^3 + 1053.4000K_t^4 - 562.69K_t^5$ 0.3000
41	Karatasou et al. [44]	Athens, Greece	TM	1996–1998	$K_t \leq 0.78$ $K_t > 0.78$	$0.9995 - 0.0500K_t - 2.4156K_t^2 + 1.4926K_t^3$ 0.2000
42	Tsubo and Walker [45]	Southern Africa	AR	2000	–	$0.613 - 0.334K_t + 0.121K_t^2$
43	Tsubo and Walker [45]	Southern Africa	AR	2000	$K_t < 0.140$ $0.140 \leq K_t \leq 0.794$ $K_t > 0.794$	0.907 0.138 0.907
44	Tsubo and Walker [45]	Southern Africa	AR	2000	$K_t < 0.140$ $0.140 \leq K_t \leq 0.794$ $K_t > 0.793$	0.907 $1.063 - 1.114K_t$ 0.180
45	Soares et al. [46]	São Paulo, Brazil	TM	1998–2001	–	$0.90 + 1.10K_t - 4.50K_t^2 + 0.01K_t^3 + 3.14K_t^4$
46	Mondol et al. [47]	Ballymena, Northern Ireland	TM	21 months of data	$K_t \leq 0.20$ $K_t > 0.20$	0.9800 $0.5836 + 3.6259K_t - 10.1710K_t^2 + 6.3380K_t^3$
47	Jacovides et al. [48]	Athalassa, Cyprus	AR	1998–2002	$K_t \leq 0.10$ $0.10 < K_t \leq 0.80$ $K_t > 0.80$	0.987 $0.940 + 0.937K_t - 5.010K_t^2 + 3.320K_t^3$ 0.177
48	Elminir et al. [49]	Aswan, Egypt	AR	1999–2001	$K_t \leq 0.22$ $0.22 < K_t \leq 0.80$ $K_t > 0.80$	$0.653 - 1.728K_t$ $0.724 - 1.821K_t + 8.221K_t^2 - 16.370K_t^3 + 9.845K_t^4$ 0.217
49	Elminir et al. [49]	Cairo, Egypt	AR	2003	$K_t \leq 0.22$ $0.22 < K_t \leq 0.80$ $K_t > 0.80$	$0.793 - 2.198K_t$ $1.341 - 9.566K_t + 32.200K_t^2 - 47.909K_t^3 + 25.419K_t^4$ 0.131
50	Elminir et al. [49]	South-Valley, Egypt	AR	2003	$K_t \leq 0.22$ $0.22 < K_t \leq 0.80$ $K_t > 0.80$	$0.8526 - 1.7780K_t$ $0.8140 - 1.1060K_t + 0.3660K_t^2 - 0.9970K_t^3 + 1.2210K_t^4$ 0.213
51	Boland et al. [50]	Adelaide, Australia	AR	–	–	$1.0/[1.0 + \exp(-5.83 + 9.87K_t)]$
52	Boland et al. [50]	Bracknell, England	TM	–	–	$1.0/[1.0 + \exp(-4.38 + 6.62K_t)]$
53	Boland et al. [50]	Darwin, Australia	TM	–	–	$1.0/[1.0 + \exp(-4.53 + 8.05K_t)]$
54	Boland et al. [50]	Lisbon, Portugal	TM	–	–	$1.0/[1.0 + \exp(-4.80 + 7.98K_t)]$
55	Boland et al. [50]	Macau, China	TM	–	–	$1.0/[1.0 + \exp(-4.87 + 8.12K_t)]$

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Table 1 (continued)

Model	Reference	Location	Climate	Data period	Constrains	$K_d$
56	Boland et al. [50]	Maputo, Mozambique	AR	–	–	$1.0/[1.0 + \exp(-5.18 + 8.80K_t)]$
57	Boland et al. [50]	Uccle, Belgium	TM	–	–	$1.0/[1.0 + \exp(-4.94 + 8.66K_t)]$
58	Boland et al. [50]	Multi-location average	Various climates	–	–	$1.0/[1.0 + \exp(-4.94 + 8.30K_t)]$
59	Boland and Ridley [51]	Multi-locations worldwide	Various climates	–	–	$1.0/[1.0 + \exp(-5.00 + 8.60K_t)]$
60	Furlan and Oliveira [52]	São Paulo, Brazil	TM	2002	$K_t \leq 0.228$	0.961
61	Mondol et al. [53]	Aldergrove, Northern Ireland	TM	1989–1998	$K_t > 0.228$	$1.337 - 1.650K_t$
					$K_t \leq 0.20$	0.9800
62	Moreno et al. [54]	Seville, Spain	TM	2000–2008	$0.20 < K_t \leq 0.70$	$0.6109 + 3.6259K_t - 10.1710K_t^2 + 6.3380K_t^3$
					$K_t > 0.70$	$0.6720 - 0.4740K_t$
					$K_t \leq 0.27$	0.9930
63	Pagola et al. [55]	3 locations in Spain	TM	2005–2008	$0.27 < K_t \leq 0.82$	$1.4946 - 1.7899K_t$
					$K_t > 0.82$	0.0450
					$K_t \leq 0.35$	$0.9818 - 0.5870K_t$
64	Pagola et al. [55]	3 locations in Spain	TM	2005–2008	$0.35 < K_t \leq 0.75$	$1.2056 - 1.3240K_t$
					$K_t > 0.75$	0.2552
					$K_t \leq 0.22$	$0.9522 - 0.3119K_t$
65	Posadillo and Lopez Luque [56]	Córdoba, Spain	TM	1993–2002	$0.22 < K_t \leq 0.80$	$0.6059 + 2.9877K_t - 10.5675K_t^2 + 10.1833K_t^3 - 3.0475K_t^4$
					$K_t > 0.80$	0.3209
66	Posadillo and Lopez Luque [56]	Córdoba, Spain	TM	1993–2002	–	$K_t(1.17 - 1.381K_t)$
67	Janjai et al. [57]	Chiang Mai, Thailand	TR	1995–2006	–	$-0.00829 + 1.16300K_t + 0.43300K_t^2 - 5.83900K_t^3 + 4.64880K_t^4$
68	Janjai et al. [57]	Nakhon Pathom, Thailand	TR	1995–2006	–	$0.9429 - 0.3707K_t + 6.4927K_t^2 - 30.3560K_t^3 + 39.1626K_t^4 - 15.4850K_t^5$
69	Janjai et al. [57]	Songkhla, Thailand	TR	1995–2006	–	$0.7699 + 2.3552K_t - 8.1480K_t^2 + 5.3811K_t^3$
70	Janjai et al. [57]	Ubon Ratchathani, Thailand	TR	1995–2006	–	$0.869 + 1.559K_t - 11.176K_t^2 + 26.143K_t^3 - 38.302K_t^4 + 31.799K_t^5 - 10.602K_t^6$
71	Ruiz-Arias et al. [27]	Albacete, Spain	TM	2002–2006	–	$0.846 + 1.841K_t - 13.425K_t^2 + 42.888K_t^3 - 85.804K_t^4 + 84.476K_t^5 - 30.637K_t^6$
72	Ruiz-Arias et al. [27]	Boulder, USA	TM	1961–1990	–	$0.086 + 0.880e^{-\exp(-3.877+6.138K_t)}$
73	Ruiz-Arias et al. [27]	Dresden, Germany	TM	1981–1990	–	$0.967 - 1.024e^{-\exp(2.473-5.324K_t)}$
74	Ruiz-Arias et al. [27]	Pittsburgh, USA	TM	1961–1990	–	$0.140 + 0.962e^{-\exp(-1.976+4.067K_t)}$
75	Ruiz-Arias et al. [27]	Savannah, USA	TM	1961–1990	–	$1.001 - 1.000e^{-\exp(2.450-5.048K_t)}$
76	Ruiz-Arias et al. [27]	Talkeetna, USA	HA	1961–1990	–	$0.988 - 1.000e^{-\exp(2.456-5.172K_t)}$
77	Ruiz-Arias et al. [27]	Tucson, USA	AR	1961–1990	–	$0.985 - 0.962e^{-\exp(2.655-6.003K_t)}$
78	Ruiz-Arias et al. [27]	7 locations in Europe and USA	Various climates	1961–2006 (Various data periods)	–	$0.988 - 1.073e^{-\exp(2.298-4.909K_t)}$
79	Torres et al. [58]	Pamplona, Spain	TM	2006–2008	$K_t \leq 0.24$	$0.952 - 1.041e^{-\exp(2.3-4.702K_t)}$
					$0.24 < K_t < 0.75$	$1.0058 - 0.2195K_t$
					$K_t \geq 0.75$	$1.3264 - 1.5120K_t$
80	Torres et al. [58]	Pamplona, Spain	TM	2006–2008	$K_t \leq 0.22$	0.1923
					$0.22 < K_t < 0.75$	$0.9920 - 0.0980K_t$
					$K_t \geq 0.75$	$1.2158 - 1.0467K_t - 0.4480K_t^2$
81	Torres et al. [58]	Pamplona, Spain	TM	2006–2008	$K_t \leq 0.22$	0.1787
					$0.24 < K_t \leq 0.755$	$0.9923 - 0.0980K_t$
					$K_t > 0.755$	$1.1459 - 0.5612K_t - 1.4952K_t^2 + 0.7103K_t^3$
82	Torres et al. [58]	Pamplona, Spain	TM	2006–2008	$K_t \leq 0.225$	0.1755
					$0.225 < K_t \leq 0.755$	$0.9943 - 0.1165K_t$
					$K_t > 0.755$	$1.4101 - 2.9918K_t + 6.4599K_t^2 - 10.3290K_t^3 + 5.5140K_t^4$
83	Chikh et al. [59]	Alger, Algeria	AR	1992	$K_t \leq 0.175$	0.1800
					$0.175 < K_t \leq 0.87$	$1.0 - 0.232K_t$
					$K_t > 0.87$	$1.170 - 1.230K_t$
84	Chikh et al. [59]	Bechar, Algeria	AR	1990–1992	$K_t \leq 0.175$	0.203
					$0.175 < K_t \leq 0.87$	$1.0 - 0.3000K_t$
					$K_t > 0.87$	$1.1370 - 1.0770K_t$
85	Chikh et al. [59]	Tamanrasset, Algeria	AR	1990–1992	$K_t \leq 0.175$	0.2043
					$0.175 < K_t \leq 0.87$	$1.0 - 0.640K_t$
					$K_t > 0.87$	$1.052 - 0.935K_t$
86	Sanchez et al. [60]	Badajoz, Spain	TM	2009–2010	$K_t < 0.30$	0.240
					$0.30 \leq K_t \leq 0.75$	0.78
					$K_t > 0.75$	$1.23 - 1.43K_t$
						0.13

(continued on next page)

Table 1 (continued)

Model	Reference	Location	Climate	Data period	Constrains	$K_d$
87	Lee et al. [61]	South Korea	TM	1986–2005	$K_t \leq 0.20$ $K_t > 0.20$	0.9200 $0.6910 + 2.4306K_t - 7.3371K_t^2 + 4.7002K_t^3$
88	Yao et al. [62]	Shanghai, China	TM	2012	$K_t \leq 0.30$ $0.30 < K_t \leq 0.80$ $K_t > 0.80$	0.9381 – 0.1481 $K_t$ 1.5197 – 1.5340 $K_t$ 0.2700
89	Yao et al. [62]	Shanghai, China	TM	2012	–	$0.8142 + 2.0792K_t - 6.1439K_t^2 + 3.4707K_t^3$
90	Yao et al. [62]	Shanghai, China	TM	2012	$K_t \leq 0.20$ $0.20 < K_t \leq 0.80$ $K_t > 0.80$	$0.8755 + 1.3991K_t - 4.9285K_t^2$ $1.1209 - 2.1699K_t + 11.0600K_t^2 - 22.3550K_t^3 + 12.8630K_t^4$ 0.2700
91	Yao et al. [62]	Shanghai, China	TM	2012	–	$0.2421 + 0.7202/[1 + \exp\{(K_t - 0.6203)/(0.0749)\}]$
92	Tapakis et al. [63]	Athalassa, Cyprus	AR	2001–2010	$K_t < 0.10$ $0.10 \leq K_t \leq 0.78$ $K_t > 0.78$	$0.9100 + 2.4993K_t - 18.8580K_t^2$ $0.9605 + 0.4482K_t - 2.0011K_t^2 - 1.5581K_t^3 + 2.0080K_t^4$ $- 2.4518 + 3.3014K_t$
93	Abreu et al. [64]	Évora, Portugal	TM	2015–2016	–	$[1.0 + (1.502 - 1.820K_t)^{-48.589}]^{-1.0/48.589}$
94	Marques Filho et al. [65]	Rio de Janeiro, Brazil	TM	2011–2014	–	$1.0/[1.0 + \exp(-4.90 + 8.78K_t)]$
95	Marques Filho et al. [65]	Rio de Janeiro, Brazil	TM	2011–2014	–	$0.13 + 0.86/[1.0 + \exp(-6.29 + 12.26K_t)]$
96	Paulescu and Blaga [66]	Timisoara, Romania	TM	2009–2010	$K_t < 0.247$ $K_t \geq 0.247$	$0.936 + 0.194K_t$ $1.436 - 1.824K_t$
97	Abal et al. [67]	Montevideo, Uruguay	TM	2011–2013	$K_t < 0.20$ $0.20 \leq K_t \leq 0.85$ $K_t > 0.85$	1.0 $0.50 + 5.92K_t - 22.22K_t^2 + 29.51K_t^3 - 19.54K_t^4 + 6.09K_t^5$ 0.10
98	Abal et al. [67]	Salto, Uruguay	TM	1998–2003	$K_t < 0.20$ $0.20 \leq K_t \leq 0.85$ $K_t > 0.85$	1.0 $0.72 + 2.80K_t - 6.62K_t^2 - 4.66K_t^3 + 14.13K_t^4 - 6.20K_t^5$ 0.09
99	Abal et al. [67]	Luján, Uruguay	TM	2011–2012	$K_t < 0.20$ $0.20 \leq K_t \leq 0.85$ $K_t > 0.85$	1.0 $0.80 + 1.97K_t - 3.93K_t^2 - 5.97K_t^3 + 10.96K_t^4 - 3.56K_t^5$ 0.11
100	Abal et al. [67]	Artigas, Uruguay	TM	2014–2015	$K_t < 0.20$ $0.20 \leq K_t \leq 0.85$ $K_t > 0.85$	1.0 $0.86 + 0.87K_t + 3.53K_t^2 - 28.43K_t^3 + 39.51K_t^4 - 16.21K_t^5$ 0.11
101	Abal et al. [67]	Treinta y Tres, Uruguay	TM	2014–2015	$K_t < 0.20$ $0.20 \leq K_t \leq 0.85$ $K_t > 0.85$	1.0 $1.04 - 1.45K_t + 13.21K_t^2 - 43.80K_t^3 + 48.79K_t^4 - 17.60K_t^5$ 0.12
102	Abal et al. [67]	Uruguay	TM	1998–2015 (Various data periods)	$K_t < 0.20$ $0.20 \leq K_t \leq 0.85$ $K_t > 0.85$	1.0 $0.77 + 2.16K_t - 3.91K_t^2 - 9.02K_t^3 + 17.00K_t^4 - 6.79K_t^5$ 0.10
103	Abal et al. [67]	Montevideo, Uruguay	TM	2011–2013	$K_t < 0.35$ $0.35 \leq K_t \leq 0.75$ $K_t > 0.75$	$1.0 - 0.40K_t$ $1.51 - 1.86K_t$ 0.12
104	Abal et al. [67]	Salto, Uruguay	TM	1998–2003	$K_t < 0.35$ $0.35 \leq K_t \leq 0.75$ $K_t > 0.75$	$1.0 - 0.29K_t$ $1.60 - 2.00K_t$ 0.10
105	Abal et al. [67]	Luján, Uruguay	TM	2011–2012	$K_t < 0.35$ $0.35 \leq K_t \leq 0.75$ $K_t > 0.75$	$1.0 - 0.24K_t$ $1.60 - 1.95K_t$ 0.14
106	Abal et al. [67]	Artigas, Uruguay	TM	2014–2015	$K_t < 0.35$ $0.35 \leq K_t \leq 0.75$ $K_t > 0.75$	$1.0 - 0.33K_t$ $1.56 - 1.93K_t$ 0.11
107	Abal et al. [67]	Treinta y Tres, Uruguay	TM	2014–2015	$K_t < 0.35$ $0.35 \leq K_t \leq 0.75$ $K_t > 0.75$	$1.0 - 0.19K_t$ $1.63 - 1.99K_t$ 0.14
108	Abal et al. [67]	Uruguay	TM	1998–2015 (Various data periods)	$K_t < 0.35$ $0.35 \leq K_t \leq 0.75$ $K_t > 0.75$	$1.0 - 0.28K_t$ $1.59 - 1.96K_t$ 0.12
109	Abal et al. [67]	Montevideo, Uruguay	TM	2011–2013	$K_t < 0.22$ $0.22 \leq K_t \leq 0.80$ $K_t > 0.80$	$1.0 - 0.24K_t$ $0.70 + 2.63K_t - 7.38K_t^2 + 1.86K_t^3 + 2.67K_t^4$ 0.13
110	Abal et al. [67]	Salto, Uruguay	TM	1998–2003	$K_t < 0.22$ $0.22 \leq K_t \leq 0.80$ $K_t > 0.80$	1.0 $0.38 + 6.54K_t - 21.25K_t^2 + 21.37K_t^3 - 6.99K_t^4$ 0.09
111	Abal et al. [67]	Luján, Uruguay	TM	2011–2012	$K_t < 0.22$ $0.22 \leq K_t \leq 0.80$ $K_t > 0.80$	$1.0 - 0.06K_t$ $0.62 + 3.70K_t - 10.83K_t^2 + 7.00K_t^3 - 0.30K_t^4$ 0.12
112	Abal et al. [67]	Artigas, Uruguay	TM	2014–2015	$K_t < 0.22$ $0.22 \leq K_t \leq 0.80$ $K_t > 0.80$	$1.0 - 0.15K_t$ $0.68 + 2.91K_t - 7.75K_t^2 + 1.47K_t^3 + 3.24K_t^4$ 0.13

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Table 1 (continued)

Model	Reference	Location	Climate	Data period	Constrains	$K_d$
113	Abal et al. [67]	Treinta y Tres, Uruguay	TM	2014–2015	$K_t < 0.22$ $0.22 \leq K_t \leq 0.80$ $K_t > 0.80$	$1.0 - 0.10K_t$ $0.85 + 0.98K_t - 0.06K_t^2 - 9.75K_t^3 + 8.62K_t^4$ 0.13
114	Abal et al. [67]	Uruguay	TM	1998–2015 (Various data periods)	$K_t < 0.22$ $0.22 \leq K_t \leq 0.80$ $K_t > 0.80$	$1.0 - 0.09K_t$ $0.60 + 3.97K_t - 11.74K_t^2 + 7.76K_t^3 - 0.28K_t^4$ 0.11
115	Abal et al. [67]	Montevideo, Uruguay	TM	2011–2013	–	$0.95 - 0.97e^{-\exp(2.96-6.07K_t)}$
116	Abal et al. [67]	Salto, Uruguay	TM	1998–2003	–	$1.00 - 1.07e^{-\exp(2.82-5.82K_t)}$
117	Abal et al. [67]	Luján, Uruguay	TM	2011–2012	–	$0.98 - 1.05e^{-\exp(2.96-5.75K_t)}$
118	Abal et al. [67]	Artigas, Uruguay	TM	2014–2015	–	$0.95 - 0.92e^{-\exp(3.57-7.32K_t)}$
119	Abal et al. [67]	Treinta y Tres, Uruguay	TM	2014–2015	–	$0.96 - 0.97e^{-\exp(3.46-6.68K_t)}$
120	Abal et al. [67]	Uruguay	TM	1998–2015 (Various data periods)	–	$0.97 - 1.01e^{-\exp(3.07-6.17K_t)}$
121	Al-Najar and Al-Khazzar [68]	Baghdad, Iraq	AR	2015	–	$1.5973 - 4.6603K_t + 5.7190K_t^2 - 2.5719K_t^3$

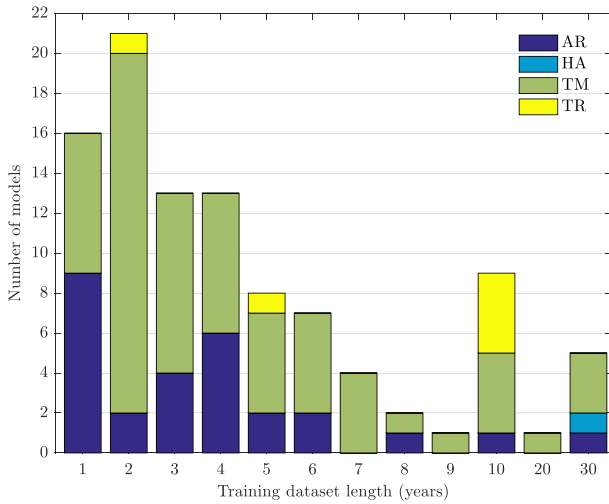


Fig. 1. Distribution of the models according to the length of the training datasets and climate zone: temperate (TM), arid (AR), tropical (TR) and high albedo (HA).

[71] and fluid flow and heat transfer optimisation when combined with the concept of the intersection of asymptotes [72,73]. In the case of modelling the diffuse fraction as a function of the clearness index, the functions used here are the physically possible limit of  $K_d = 1$  and a function  $Z$  that is the best fit for the clear sky periods. The occurrence of cloud enhancement effects in 1-min resolution data is quite frequent [1,14], then function  $Z$  was defined as the best fit to the clear sky data using a second degree polynomial in the form:

$$Z = A(K_t - 0.5)^2 + B(K_t - 0.5) + 1 \quad (2)$$

Since  $K_d$  is a concave function of  $K_t$ , the exponents used in Eq. (1) must be  $-n$  and  $-1/n$ , and thus, the final form of the model is given by:

$$K_d = \{1 + [A(K_t - 0.5)^2 + B(K_t - 0.5) + 1]^{-n}\}^{-\frac{1}{n}} \quad (3)$$

Fig. 2 shows the fitting of the model to the data of Fort Peck station (FPE), USA, as an example of the procedures implemented in this work for all the analysed stations. Red lines represent the limiting functions  $K_d = 1$  and  $Z$  (fitted to the FPE dataset), and the blue line represents the fitted model. The three parameters of the adjusted model for FPE station are also presented, as well as the period of data used.

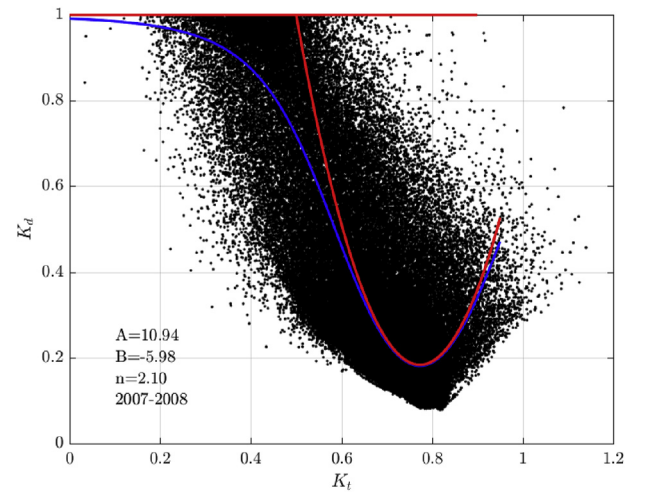


Fig. 2. Data for the FPE station (Fort Peck, USA) and representation of the limiting functions (red) and model correlation (blue). (For interpretation of the references to colour in this figure legend, the reader is referred to the Web version of this article.)

### 3.2. Test stations and data quality control

The data used in this study is from the Baseline Surface Radiation Network (BSRN) [69,74] and the Institute of Earth Sciences (IES) at the University of Évora, Portugal. The BSRN is a project of the Global Energy and Water Cycle Experiment (GEWEX) under the umbrella of the World Climate Research Programme (WCRP). The primary objective of this network is to detect changes in the radiation field at the Earth's surface which may be related to climate changes. Measurements of solar radiation in the IES station are taken likewise as in the BSRN stations: the diffuse horizontal irradiation (DHI) is measured by a Kipp&Zonnen CM6B pyranometer and shading ball attached to the sun tracker and the global horizontal irradiance is also measured by a Kipp&Zonnen Solys2 sun tracker and are properly maintained and calibrated according to the BSRN and WMO guidelines and recommendations [69,75]. Table 2 shows detailed information on the stations used in this study: location, climate zone, data period, number of valid data points and mean GHI of all valid measurements.

Data quality control was performed by applying the quality filters defined by Long and Shi [76] for the global horizontal irradiance (GHI). Furthermore,  $K_d$  values higher than 1 and lower than 0 were removed

**Table 2**

Information on the data of BSRN and IES stations. Acronyms: AR (Arid), HA (High albedo), TM (Temperate), and TR (Tropical).

Station	Code	Lat. (°N)	Long. (°E)	Elev. (m)	Climate	Data period	Samples	Mean GHI ( $W/m^2$ )
Alert	ALE	82.490	-62.420	127	HA	2009–2011	631284	223.84
Alice Springs	ASP	-23.798	133.888	547	AR	2007–2009	704784	561.96
Bermuda	BER	32.267	-64.667	8	TM	2006–2008	458837	497.33
Billings	BIL	36.605	-97.516	317	TM	2005–2007	631822	451.65
Bondville	BON	40.067	-88.367	213	TM	2007–2009	220046	507.54
Boulder	BOU	40.050	-105.007	1577	TM	2002–2004	481571	519.33
Brasilia	BRB	-15.601	-47.713	1023	TR	2009–2011	598116	472.22
Carpentras	CAR	44.083	5.059	100	TM	2003–2005	619585	419.19
Chesapeake Light	CLH	36.905	-75.713	37	TM	2011–2013	681269	408.60
Cener	CNR	42.816	-1.601	471	TM	2010–2012	666202	374.60
Cocos Island	COC	-12.193	96.835	6	TR	2006–2008	556163	505.99
De Aar	DAA	-30.667	23.993	1287	AR	2002–2004	592787	518.07
Darwin	DAR	-12.425	130.891	30	TR	2009–2011	699416	489.94
Concordia station	DOM	-75.100	123.383	3233	HA	2005–2007	255370	377.86
Desert Rock	DRA	36.626	-116.018	1007	AR	2007–2009	324644	596.47
Évora	EVR	38.568	-7.912	293	TM	2016–2017	199169	504.11
Eureka	EUR	79.989	-85.940	85	HA	2009–2011	654421	230.99
Fort Peck	FPE	48.317	-105.100	634	TM	2007–2009	227621	481.02
Fukuoka	FUA	33.582	130.376	3	TM	2011–2013	715368	337.10
Goodwin Creek	GCR	34.250	-89.870	98	TM	2007–2009	244471	529.24
Gobabeb	GOB	-23.561	15.042	407	AR	2012–2014	627165	596.64
Georg von Neumayer	GVN	-70.650	-8.250	42	HA	2011–2013	509631	316.63
Ilorin	ILO	8.533	4.567	350	TR	1995,1999,2000	160661	307.77
Ishigakijima	ISH	24.337	124.163	6	TM	2011–2013	710421	374.30
Izana	IZA	28.309	-16.499	2373	AR	2011–2013	680660	612.14
Kwajalein	KWA	8.720	167.731	10	TR	1998–2000	517467	544.96
Lauder	LAU	-45.045	169.689	350	TM	2005–2007	583349	399.17
Lerwick	LER	60.139	-1.185	80	TM	2004–2006	586958	213.18
Lindenberg	LIN	52.210	14.122	125	TM	2001–2003	665675	285.77
Momote	MAN	-2.058	147.425	6	TR	2008–2010	689159	470.82
Minamitorishima	MNM	24.288	153.983	7	TM	2011–2013	727974	470.67
Nauru Island	NAU	-0.521	166.917	7	TR	2005–2007	649304	513.96
Ny-Alesund	NYA	78.925	11.930	11	HA	2007–2009	619576	187.06
Palaiseau	PAL	48.713	2.208	156	TM	2009–2011	701389	302.24
Payerne	PAY	46.815	6.944	491	TM	2008–2010	573646	349.50
Rock Springs	PSU	40.720	-77.933	376	TM	2007–2009	203868	471.41
Petrolina	PTR	-9.068	-40.319	387	TR	2007–2009	149097	566.04
Regina	REG	50.205	-104.713	578	TM	2009–2011	620001	365.50
Sapporo	SAP	43.060	141.328	17	AR	2011–2013	722699	305.91
Sede Boqer	SBO	30.860	34.779	480	TM	2009–2011	662564	542.45
Sonnblick	SON	47.054	12.958	3109	HA	2013–2015	462080	371.23
Solar Village	SOV	24.907	46.397	768	AR	2000–2002	717011	564.91
Sioux Falls	SXF	43.730	-96.620	473	TM	2007–2009	228600	475.41
Tamanrasset	TAM	22.790	5.529	1385	AR	2006–2008	644908	596.17
Tateno	TAT	36.050	140.133	25	TM	2008–2010	683402	334.39
Tiksi	TIK	71.586	128.919	48	HA	2011–2013	617716	211.72
Toravere	TOR	58.254	26.462	70	TM	2010–2012	649189	245.32
Xianghe	XIA	39.754	116.962	32	TM	2008–2010	561393	381.54

for the fitting of the model parameters  $A$  and  $B$ , since measurements of diffuse irradiance higher than global irradiance are very dubious. However, instrumental errors can occur, and therefore, the  $K_d$  maximum value was set to 1.2 when determining the parameter  $n$ . Finally, measurements taken when the zenith angle was higher than  $85^\circ$  were also removed due to disturbances caused mainly by the horizon line and also due to instrumental and modelling accuracy issues in that case [14]. Since 1-min data was used in this work, the extraterrestrial irradiance on a horizontal surface,  $E_{oh}$ , that is needed to determine the  $K_t$  was simply calculated based on the solar constant ( $G_s = 1361.1 Wm^{-2}$  [77]) and the solar zenith angle, this last being calculated through the very accurate solar position algorithm developed by Reda and Andreas [78]. The data was divided into two sets: the training set with two years of data used to determine the fitting parameters of the proposed model; and the validation set with one year of data, used to validate the developed model as well as the models available in the literature (Section 2). These datasets are composed of years with high number of valid measurements, i.e., records that successfully passed the quality control procedures [76], as shown in Table 2.

### 3.3. Statistical indicators for model assessment

The developed model, as well as the models reviewed in Section 2, were evaluated using the statistical indicators described below taking the measured values as reference. Lower values indicate better model accuracy except for the mean bias error and mean percentage error, in which values closer to zero indicate a better model accuracy, and for the correlation coefficient, in which a value closer to 1 represents better model accuracy. In the following,  $H$  and  $N$  stand for minutely diffuse horizontal irradiance (DHI) and number of observations, respectively, and the subscripts  $m$ ,  $e$  and  $avg$  stand for measured, estimated and average, respectively.

#### 3.3.1. Mean bias error (MBE)

$$MBE = \frac{1}{N} \sum_{i=1}^N (H_{e,i} - H_{m,i}) \quad (4)$$



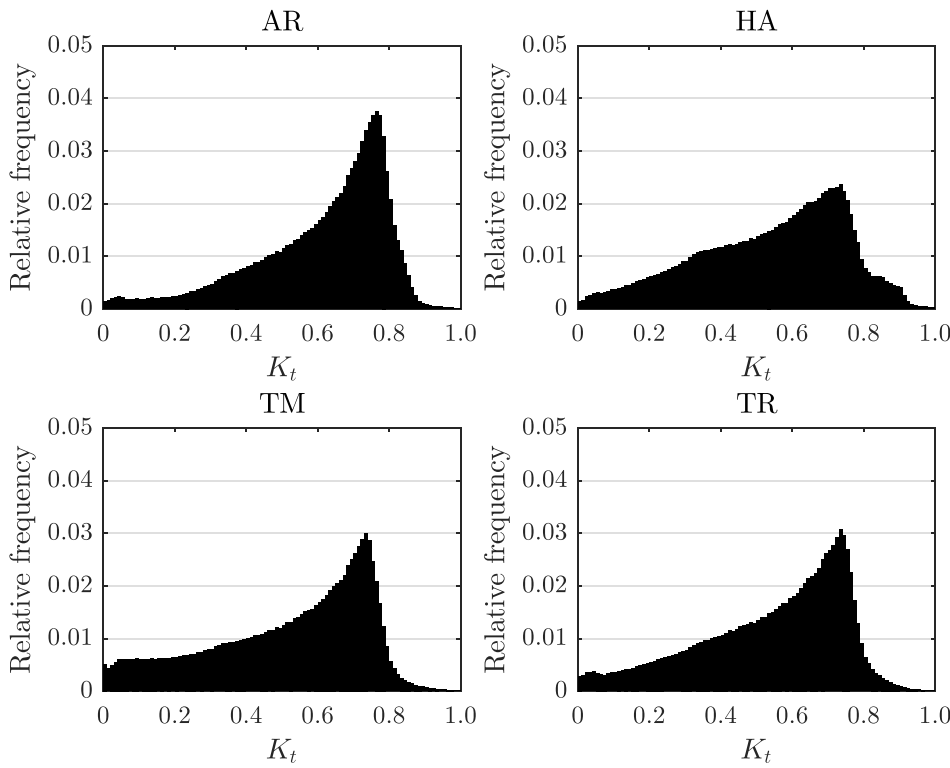


Fig. 3. Relative frequency of  $K_t$  in BSRN stations according to the climate zone.

3.3.2. Mean absolute error (MAE)

$$MAE = \frac{1}{N} \sum_{i=1}^N |H_{e,i} - H_{m,i}| \tag{5}$$

3.3.3. Root mean square error (RMSE)

$$RMSE = \left[ \frac{1}{N} \sum_{i=1}^N (H_{e,i} - H_{m,i})^2 \right]^{\frac{1}{2}} \tag{6}$$

3.3.4. Mean percentage error (MPE)

$$MPE = \frac{1}{N} \sum_{i=1}^N \left( \frac{H_{e,i} - H_{m,i}}{H_{m,i}} \right) \times 100 \tag{7}$$

3.3.5. Uncertainty at 95% (U95)

$$U95 = 1.96(SD^2 + RMSE^2)^{\frac{1}{2}} \tag{8}$$

where  $SD$  represents the standard deviation of the difference between  $H_e$  and  $H_m$ .

3.3.6. Relative root mean square error (RRMSE)

$$RRMSE = \frac{RMSE}{H_{m,avg}} \tag{9}$$

3.3.7. T-statistics (TSTAT)

$$TSTAT = \left[ \frac{(N - 1)MBE^2}{RMSE^2 - MBE^2} \right]^{\frac{1}{2}} \tag{10}$$

3.3.8. Maximum absolute relative error (erMAX)

$$erMAX = \max \left\{ \left| \frac{H_{e,i} - H_{m,i}}{H_{m,i}} \right|, i = 1, \dots, N \right\} \tag{11}$$

3.3.9. Correlation coefficient (R)

$$R = \frac{\sum_{i=1}^N (H_{e,i} - H_{e,avg})(H_{m,i} - H_{m,avg})}{\sqrt{\sum_{i=1}^N (H_{e,i} - H_{e,avg})^2 \sum_{i=1}^N (H_{m,i} - H_{m,avg})^2}} \tag{12}$$

3.3.10. Mean absolute relative error (MARE)

$$MARE = \frac{1}{N} \sum_{i=1}^N \left| \frac{H_{e,i} - H_{m,i}}{H_{m,i}} \right| \tag{13}$$

3.3.11. Global performance index (GPI)

The global performance index, firstly proposed by Behar et al. [79], then modified by Despotovic et al. [80] and used by Jamil and Akthar [23,81], is also used here to combine all the statistical indicators presented in Subsections 3.3.1 to 3.3.10. The need for using this index is due to the incapacity of those statistical indicators to, consistently, identify the best model (see Table 4). The values of the statistical indicators need to be scaled between 0 (worst performing model) and 1 (best performing model) to determine the GPI values, which otherwise would make difficult to compare the models. This normalisation also allows using the same statistical weight for all of the indicators when determining the GPI, as follows:

$$GPI = \sum_{j=1}^{10} \alpha_j (\bar{y}_j - y_{ij}) \tag{14}$$

where  $\bar{y}_j$  is the median of the scaled values of the indicator  $j$ ,  $y_{ij}$  is the scaled value of the statistical indicator  $j$  for the model  $i$ , and  $\alpha_j$  equals to 1 for all statistical indicators except  $R$ , in which  $\alpha_j$  equals to  $-1$ . The GPI also allows to combine several indicators regardless if the value of a single indicator is 0 or not, which is not possible if a simple product of

the indicators is computed, and therefore a higher GPI stands for better accuracy of a given model.

#### 4. Results and discussion

##### 4.1. Determination of model parameters and climate analysis

Fig. 3 shows the distribution of the  $K_t$  values for the climate zones considered, which is useful to identify the range of the clearness index and the clear sky occurrences (frequency). The AR climate zone presents the highest relative frequency for high values of  $K_t$ , reaching a relative frequency of 0.038 for  $K_t \approx 0.77$ , followed by the TM and TR climate zones with maximum relative frequency around 0.030 for approximately the same  $K_t$  value. The HA climate zone presents the lower values of  $K_t$  relative frequency for clear sky.

The training of the model was performed using the stations from the BSRN network, while the IES station (EVR) was used only in the validation of the model. The parameters  $A$  and  $B$  for each station were found by fitting Eq. (2) to the data in the range of  $K_t \geq 0.5$ , using the non-linear least squares method. Then, the parameter  $n$  was obtained through an optimisation process in order to achieve the maximum GPI value for the entire range of the data. One also investigated the existence of a possible relationship between these parameters and the elevation of the stations according to the climate zone, as shown in Fig. 4. However no conclusions can be drawn on the existence of any clear dependence, and therefore, no traditional fitting equations using the parameters of the model and the elevation of the stations were able to present an acceptable coefficient of determination.

To develop a model based only on the  $K_t$  as predictor for each climate region, the mean value and standard deviation of these parameters were calculated for the four zones considered, and the stations in which at least one of the three parameters were out of the confidence interval defined by the mean  $\pm$  standard deviation were excluded from

**Table 3**  
Parameters of the developed model according to the climate zone: Arid (AR), High Albedo (HA), Temperate (TM) and Tropical (TR).

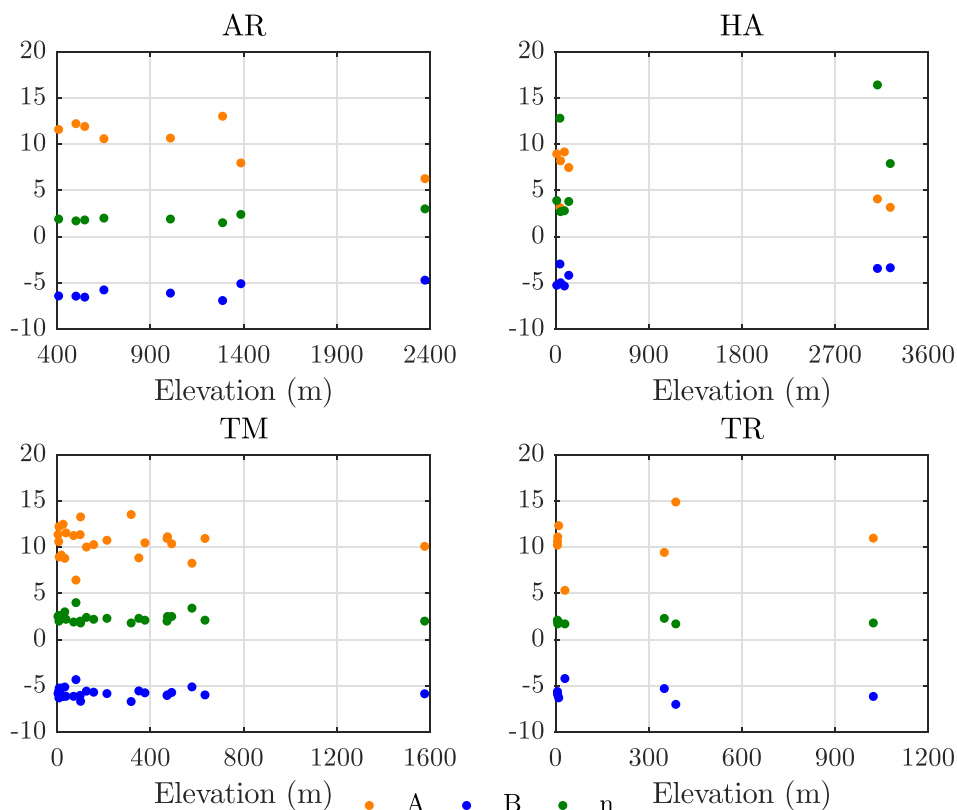
Parameters	Climate Zone			
	AR	HA	TM	TR
A	11.39	7.83	10.79	11.59
B	-6.25	-4.59	-5.87	-6.14
n	1.86	3.25	2.24	1.87

this calculation. After this procedure, the values inside this imposed range are averaged in order to obtain the mean values of  $A$ ,  $B$  and  $n$  for each climate zone, as shown in Table 3. These parameters were used for the model performance assessment as presented in the following section.

##### 4.2. Performance assessment

The statistical tools presented in Section 3.3 were used to assess the performance of the models from the literature as well as of the model developed in this work using measurements from the EVR station and the datasets from the BSRN, as presented in Table 2. The performance assessment was carried out using the corresponding set of parameters according to the climate zone of the 48 radiometric stations analysed in this study. The performance assessment is presented in detail for the EVR station as an example of both the methodology used in this study and as a completely independent assessment since data from this station was not included in the determination of the model parameters. Table 4 presents the results of the statistical analysis of the selected models using the indicators shown in Section 3.3 for the EVR station. The bold font indicates the optimal values of the statistical indicators, i.e., it indicates the best model according to each statistical indicator.

The statistical evaluation regarding the accuracy of the models



**Fig. 4.** Variation of the parameters  $A$ ,  $B$  and  $n$  according to the elevation of the stations and climate zone: Arid (AR), High Albedo (HA), Temperate (TM) and Tropical (TR).

**Table 4**  
Statistical analysis of the selected models for the EVR station.

Model	MBE ( $W\ m^{-2}$ )	RMSE ( $W\ m^{-2}$ )	TSTAT	MPE (%)	U95 ( $W\ m^{-2}$ )	erMAX	RRMSE	MAE ( $W\ m^{-2}$ )	R	MARE	GPI
1	-40.31	111.13	36.42	-6.42	297.74	2.00	0.86	75.41	0.19	0.55	0.22
2	-29.87	102.79	28.42	3.82	278.84	2.11	0.79	71.26	0.30	0.54	0.64
3	-46.36	112.67	42.25	-12.76	298.80	1.79	0.87	75.05	0.19	0.53	0.13
4	-74.08	124.44	69.34	-41.84	312.87	<b>0.89</b>	0.96	79.87	0.19	0.50	-0.46
5	-79.12	127.30	74.25	-47.13	316.96	0.90	0.98	82.42	0.19	0.52	-0.63
6	-74.08	124.44	69.34	-41.84	312.87	<b>0.89</b>	0.96	79.87	0.19	0.50	-0.46
7	-99.29	140.53	93.43	-68.28	337.45	0.94	1.08	99.29	0.19	0.68	-1.46
8	-79.12	127.30	74.25	-47.13	316.96	0.90	0.98	82.42	0.19	0.52	-0.63
9	-53.92	115.14	49.60	-20.69	301.16	1.54	0.89	75.16	0.19	0.51	0.00
10	-79.12	127.30	74.25	-47.13	316.96	0.90	0.98	82.42	0.19	0.52	-0.63
11	-69.04	121.78	64.41	-36.56	309.23	1.03	0.94	77.95	0.19	0.49	-0.32
12	-58.96	117.11	54.53	-25.98	303.36	1.37	0.90	75.67	0.19	0.50	-0.10
13	-94.25	136.99	88.73	-62.99	331.75	0.93	1.06	94.31	0.19	0.63	-1.24
14	-79.12	127.30	74.25	-47.13	316.96	0.90	0.98	82.42	0.19	0.52	-0.63
15	-89.21	133.59	83.95	-57.70	326.42	0.92	1.03	89.69	0.19	0.58	-1.02
16	-79.12	127.30	74.25	-47.13	316.96	0.90	0.98	82.42	0.19	0.52	-0.63
17	-21.15	109.03	18.51	13.67	299.35	2.64	0.84	78.71	0.19	0.64	0.24
18	28.34	79.71	35.61	56.30	213.84	2.98	0.62	63.06	0.68	0.69	0.91
19	-48.88	113.43	44.69	-15.41	299.45	1.71	0.88	75.01	0.19	0.52	0.09
20	-38.25	107.47	35.65	-3.12	288.30	2.10	0.83	69.93	0.24	0.49	0.45
21	1.53	112.10	1.28	37.46	310.72	3.40	0.87	85.97	0.19	0.78	-0.01
22	-16.11	109.19	13.96	18.96	301.01	2.81	0.84	80.05	0.19	0.67	0.20
23	<b>0.42</b>	82.32	<b>0.48</b>	28.96	228.18	2.45	0.64	54.46	0.59	0.51	1.66
24	-55.43	115.71	51.08	-22.28	301.76	1.49	0.89	75.28	0.19	0.50	-0.03
25	-53.41	114.96	49.11	-20.17	300.96	1.56	0.89	75.13	0.19	0.51	0.01
26	-40.31	111.13	36.42	-6.42	297.74	2.00	0.86	75.41	0.19	0.55	0.22
27	-30.22	109.49	26.88	4.15	297.65	2.34	0.85	76.76	0.19	0.59	0.29
28	3.55	112.66	2.95	39.58	312.20	3.47	0.87	86.76	0.19	0.80	-0.09
29	19.18	118.40	15.36	55.97	326.03	3.99	0.91	93.67	0.19	0.91	-0.69
30	8.09	114.07	6.65	44.34	315.79	3.62	0.88	88.64	0.19	0.83	-0.26
31	-99.29	140.53	93.43	-68.28	337.45	0.94	1.08	99.29	0.19	0.68	-1.46
32	-56.45	105.00	59.68	-26.47	269.19	1.77	0.81	60.44	0.49	0.33	0.64
33	-47.02	121.55	39.26	-12.54	324.08	2.03	0.94	82.07	0.08	0.60	-0.29
34	-129.12	164.09	119.34	-99.48	377.92	1.00	1.27	129.12	0.18	0.99	-2.82
35	14.18	77.91	17.32	42.93	214.17	2.63	0.60	58.39	0.65	0.60	1.37
36	-38.79	110.81	34.98	-4.83	297.59	2.05	0.86	75.56	0.19	0.56	0.24
37	-13.59	109.39	11.72	21.60	302.03	2.89	0.84	80.77	0.19	0.69	0.18
38	21.99	87.63	24.26	45.07	239.04	5.94	0.68	60.85	0.61	0.60	0.49
39	21.70	119.55	17.27	58.61	328.63	4.08	0.92	94.89	0.19	0.93	-0.79
40	21.70	119.55	17.27	58.61	328.63	4.08	0.92	94.89	0.19	0.93	-0.79
41	-28.71	109.34	25.47	5.74	297.82	2.39	0.84	77.04	0.19	0.60	0.28
42	91.34	162.12	63.82	130.84	412.18	6.25	1.25	131.34	0.23	1.48	-3.61
43	-59.97	117.54	55.52	-27.04	303.86	1.34	0.91	75.82	0.19	0.50	-0.12
44	-38.79	110.81	34.98	-4.83	297.59	2.05	0.86	75.56	0.19	0.56	0.24
45	4.17	75.61	5.17	28.26	209.41	4.19	0.58	52.45	0.67	0.47	1.56
46	35.49	81.81	45.07	63.51	215.83	3.10	0.63	65.77	0.69	0.74	0.64
47	-40.31	111.13	36.42	-6.42	297.74	2.00	0.86	75.41	0.19	0.55	0.22
48	-20.14	109.03	17.59	14.73	299.64	2.67	0.84	78.97	0.19	0.65	0.23
49	-63.50	119.10	58.98	-30.74	305.76	1.22	0.92	76.47	0.19	0.49	-0.20
50	-22.16	109.03	19.43	12.61	299.08	2.61	0.84	78.47	0.19	0.64	0.25
51	0.97	74.95	1.21	27.30	207.75	2.63	0.58	50.05	0.67	0.47	1.92
52	80.04	123.80	79.32	114.80	305.19	4.66	0.96	101.68	0.51	1.22	-1.94
53	-1.70	75.32	2.12	25.91	208.75	2.26	0.58	51.50	0.67	0.47	1.96
54	25.21	82.06	30.22	54.78	222.02	2.88	0.63	64.16	0.64	0.69	0.89
55	22.66	80.81	27.34	51.93	219.55	2.85	0.62	62.57	0.64	0.67	1.01
56	8.51	75.99	10.54	36.20	209.99	2.64	0.59	54.50	0.67	0.54	1.62
57	-2.72	74.91	3.40	24.30	207.57	2.33	0.58	50.11	0.67	0.46	1.97
58	17.72	78.78	21.61	46.49	215.59	2.77	0.61	59.71	0.65	0.62	1.23
59	5.25	75.53	6.53	32.87	209.11	2.52	0.58	53.32	0.67	0.52	1.75
60	-48.27	108.42	46.53	-19.98	285.25	2.48	0.84	57.67	0.42	0.36	0.56
61	38.92	113.31	34.23	75.18	304.68	4.23	0.87	92.15	0.31	0.97	-0.91
62	-106.85	146.12	100.33	-76.21	346.67	0.96	1.13	106.85	0.19	0.76	-1.80
63	-0.89	111.49	0.74	34.92	309.04	3.32	0.86	85.04	0.19	0.77	0.06
64	32.23	124.95	24.99	69.66	340.53	4.43	0.96	100.26	0.19	1.01	-1.23
65	-72.77	123.26	68.46	-39.64	310.46	2.34	0.95	74.37	0.23	0.42	-0.50
66	-57.44	104.79	61.34	-28.93	267.76	2.56	0.81	65.17	0.51	0.39	0.43
67	17.89	77.44	22.22	44.70	211.78	2.80	0.60	58.85	0.67	0.60	1.30
68	17.38	74.36	22.50	44.68	203.28	2.59	0.57	57.34	0.70	0.60	1.44
69	54.08	108.58	53.76	86.97	281.68	4.31	0.84	87.56	0.47	0.99	-0.92
70	10.66	75.44	13.35	38.81	208.07	6.90	0.58	56.61	0.68	0.56	0.86
71	8.27	74.35	10.48	36.44	205.45	2.64	0.57	55.35	0.69	0.55	1.66
72	-24.73	80.84	30.08	2.11	218.76	1.88	0.62	47.74	0.66	0.33	1.87
73	-7.55	78.20	9.08	21.51	216.24	2.09	0.60	55.36	0.66	0.47	1.84
74	18.79	80.38	22.50	48.26	219.75	2.67	0.62	61.71	0.64	0.64	1.15

(continued on next page)

Table 4 (continued)

Model	MBE ( $W m^{-2}$ )	RMSE ( $W m^{-2}$ )	TSTAT	MPE (%)	U95 ( $W m^{-2}$ )	erMAX	RRMSE	MAE ( $W m^{-2}$ )	R	MARE	GPI
75	3.70	76.96	4.50	32.28	213.19	2.17	0.59	55.20	0.65	0.52	1.78
76	-16.42	77.36	20.33	10.66	211.99	1.89	0.60	49.07	0.68	0.38	1.97
77	-20.39	80.17	24.61	6.91	218.59	1.91	0.62	48.59	0.65	0.36	1.87
78	-9.96	78.53	11.97	18.35	216.80	1.92	0.61	51.70	0.64	0.43	1.88
79	-32.59	109.77	29.10	1.67	297.48	2.26	0.85	76.37	0.19	0.58	0.29
80	-39.45	110.95	35.60	-5.52	297.65	2.03	0.86	75.49	0.19	0.55	0.23
81	-41.06	111.30	37.15	-7.21	297.83	1.97	0.86	75.35	0.19	0.55	0.21
82	-38.79	110.81	34.98	-4.83	297.59	2.05	0.86	75.56	0.19	0.56	0.24
83	-27.20	109.23	24.06	7.33	298.03	2.44	0.84	77.34	0.19	0.61	0.27
84	-26.54	109.18	23.46	8.01	298.14	2.46	0.84	77.48	0.19	0.61	0.27
85	-8.55	110.00	7.29	26.89	304.45	3.06	0.85	82.35	0.19	0.72	0.14
86	-64.00	119.33	59.47	-31.27	306.06	1.20	0.92	76.58	0.19	0.49	-0.21
87	84.34	128.86	81.01	116.82	316.63	5.10	0.99	103.80	0.50	1.24	-2.17
88	6.58	113.58	5.43	42.75	314.55	3.57	0.88	88.00	0.19	0.82	-0.20
89	92.63	132.94	90.92	126.87	320.66	5.13	1.03	108.04	0.51	1.32	-2.49
90	88.67	123.26	96.91	119.65	294.17	4.67	0.95	100.93	0.61	1.24	-2.00
91	87.44	121.89	96.37	119.29	291.17	4.63	0.94	100.66	0.61	1.24	-1.97
92	-179.35	287.33	74.77	-141.81	714.66	5.68	2.22	202.48	-0.30	1.59	-6.35
93	-18.41	72.93	24.41	5.27	198.91	2.15	0.56	44.95	0.72	0.36	2.08
94	-11.85	75.52	14.87	14.47	208.03	2.23	0.58	47.25	0.68	0.39	1.98
95	-5.82	71.58	7.64	19.72	198.07	2.30	0.55	47.76	0.72	0.41	2.12
96	-60.98	121.20	54.48	-33.86	313.97	2.74	0.94	67.90	0.38	0.46	-0.10
97	-79.12	127.30	74.25	-47.13	316.96	0.90	0.98	82.42	0.19	0.52	-0.63
98	-84.16	130.36	79.13	-52.42	321.48	0.91	1.01	85.68	0.19	0.55	-0.82
99	-74.08	124.44	69.34	-41.84	312.87	0.89	0.96	79.87	0.19	0.50	-0.46
100	-74.08	124.44	69.34	-41.84	312.87	0.89	0.96	79.87	0.19	0.50	-0.46
101	-69.04	121.78	64.41	-36.56	309.23	1.03	0.94	77.95	0.19	0.49	-0.32
102	-79.12	127.30	74.25	-47.13	316.96	0.90	0.98	82.42	0.19	0.52	-0.63
103	-69.04	121.78	64.41	-36.56	309.23	1.03	0.94	77.95	0.19	0.49	-0.32
104	-79.12	127.30	74.25	-47.13	316.96	0.90	0.98	82.42	0.19	0.52	-0.63
105	-58.96	117.11	54.53	-25.98	303.36	1.37	0.90	75.67	0.19	0.50	-0.10
106	-74.08	124.44	69.34	-41.84	312.87	0.89	0.96	79.87	0.19	0.50	-0.46
107	-58.96	117.11	54.53	-25.98	303.36	1.37	0.90	75.67	0.19	0.50	-0.10
108	-69.04	121.78	64.41	-36.56	309.23	1.03	0.94	77.95	0.19	0.49	-0.32
109	-64.00	119.33	59.47	-31.27	306.06	1.20	0.92	76.58	0.19	0.49	-0.21
110	-84.16	130.36	79.13	-52.42	321.48	0.91	1.01	85.68	0.19	0.55	-0.82
111	-69.04	121.78	64.41	-36.56	309.23	1.03	0.94	77.95	0.19	0.49	-0.32
112	-64.00	119.33	59.47	-31.27	306.06	1.20	0.92	76.58	0.19	0.49	-0.21
113	-64.00	119.33	59.47	-31.27	306.06	1.20	0.92	76.58	0.19	0.49	-0.21
114	-74.08	124.44	69.34	-41.84	312.87	0.89	0.96	79.87	0.19	0.50	-0.46
115	-15.14	76.83	18.81	11.50	210.89	2.11	0.59	47.82	0.67	0.38	1.95
116	-24.37	80.74	29.63	1.31	218.65	2.19	0.62	45.81	0.65	0.32	1.84
117	-6.69	76.75	8.19	20.13	212.34	2.34	0.59	49.23	0.66	0.43	1.91
118	-14.95	75.02	19.03	10.74	205.86	2.27	0.58	46.33	0.69	0.36	2.01
119	-2.52	74.47	3.16	23.97	206.36	2.44	0.57	49.39	0.68	0.45	1.98
120	-16.08	77.31	19.90	10.01	211.97	2.25	0.60	46.70	0.67	0.36	1.93
121	1.42	92.87	1.44	34.45	257.41	2.85	0.72	69.35	0.42	0.62	0.95
This work	25.84	70.80	36.68	47.87	189.61	2.59	0.55	53.70	0.76	0.58	1.43

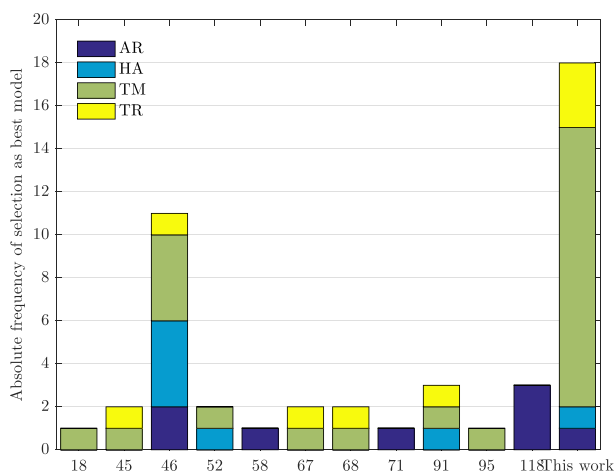


Fig. 5. Best performing models according to the climate zone: arid (AR), high albedo (HA), temperate (TM) and tropical (TR).

presented in Table 4 for EVR does not give a unanimous decision on which of the models is considered the most accurate. The model developed in this work is the best performing model if we consider the RMSE ( $70.80 W m^{-2}$ ), U95 ( $189.61 W m^{-2}$ ), RRMSE (0.55) and R (0.76) statistical indicators. On the other hand, the model 23 presented by Hollands [31], is the most accurate model regarding the MBE ( $0.42 W m^{-2}$ ) and TSTAT (0.48) indicators. Considering the values of MPE (1.31%) and MARE (0.32) of the model 116 derived by Abal et al. [67], it would be selected as the best performing model for this station. Regarding the erMAX, it does not allow identifying the single most accurate model, due to presenting the same value for several models. The analysis of these results shows the advantage of using the GPI in order to present a more concise performance evaluation, allowing the combination of several statistical indicators and providing, through a simple procedure, a result easy to understand. Therefore, the model with higher GPI for the station being analysed (EVR) is the model 95 presented by Marques Filho et al. [65] with a GPI value of 2.12, although it was not considered the best performing model according to any of the statistical indicators separately. This result is due to the scaling down of the values of the statistical indicators mentioned above in the GPI determination procedure, which allows for a fair comparison of the

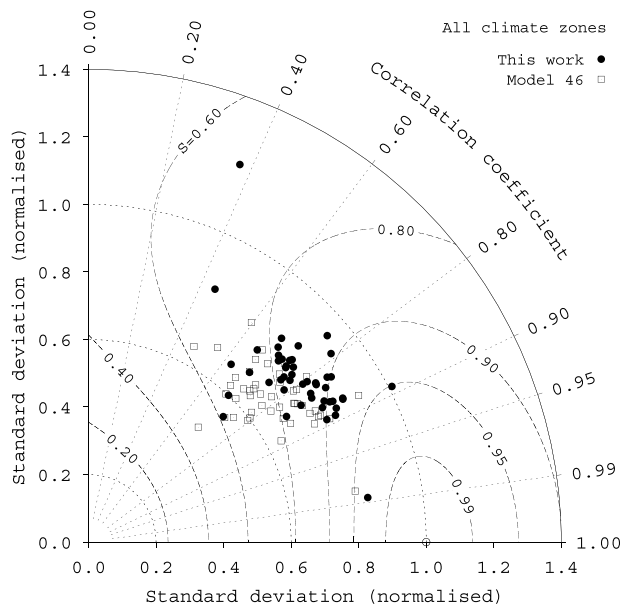


Fig. 6. Taylor diagram for the results of model 46 and the model developed in this work for all climate zones.

models under study.

This analysis was also carried out for all of the BSRN stations in order to identify the best performing model for various locations in different climate zones. Fig. 5 presents the result of this comprehensive performance assessment, where the absolute frequency of selection as the best performing model (based on the maximum GPI) is shown, also according to the climate zone of the stations.

The best overall performing model is the one presented in this work (Eq. (3) and Table 3), followed by model 46 by Mondol et al. [47]. The best performing model for the AR climate zone is the model 118 proposed by Abal et al. [67] while the best performing model for the HA climate zone is the model 46 derived by Mondol et al. [47]. The best performing model for the TM and TR climate zones is the model presented in this work (Eq. (3) and Table 3). The results presented in Fig. 5 show the advantage in deriving specific model parameters for each climate zone as suggested by Gueymard and Ruiz-Arias [14], despite the good performance of the model proposed by Mondol [47]. Only the model proposed here, and the model proposed by Mondol [47] were selected as best performing models for the four climate zones considered, showing excellent versatility regarding climate zone selection. On the other hand, the higher number of stations in the temperate (TM) zone allows for a better characterisation of the model parameters in comparison against other climate zones, where less data is available.

To further test the closeness of the two best performing models from Fig. 5 to the measured data, i.e., model 46 and the one proposed in this work, a Taylor diagram [25] is presented in Fig. 6 for all stations. This diagram helps to identify the closeness between the modelled and observed data in terms of their correlation coefficient (azimuthal position,  $R$ ) and standard deviation of model output normalised by the standard deviation of the corresponding observations (radial distance,  $nSD$ ). In this diagram, a better accuracy is achieved when the model is located close to the normalised standard deviation unit line [14], while the open circle in the coordinates ( $nSD = 1$ ,  $R = 1$ ) represents the point of perfect match between model and observations. It is worth to mention that the distance between the model representation and this point is the so-called centred pattern Root Mean Square Difference, which accounts for the mean values of both data series, and is also a measure of the model closeness to the observed data (not shown). Instead, a more elaborated skill score proposed by Taylor [25] was determined as a function of the correlation coefficient ( $R$ ) and the normalised standard

deviation ( $nSD$ ) in the form  $S = 4(1 + R)/[(nSD + 1/nSD)^2(1 + R_0)]$  with  $R_0 = \max(R)$ , which is represented in Fig. 6 by the  $S$  isolines. This definition allows guaranteeing that the skillfulness of the model increases when  $nSD \rightarrow 1$ , which is not assured by the centred pattern RMSE for lower values of  $R$ . Best model accuracy is achieved when skill score values tend to unity. In this way, the best performing model according to Fig. 6 is the model proposed in this work, with a skill score higher than 0.80 for the majority of the stations and with  $nSD$  values in general higher than those of model 46. This conclusion is in agreement with the results presented in Fig. 5, showing that the proposed model has best accuracy, using two distinct statistical analysis.

## 5. Conclusions

In this work is presented a newly developed model based on minute resolution data and climate zone classification. A review of the literature on models that use hourly and sub-hourly  $K_t$  values to compute  $K_d$  was carried out. An extensive statistical analysis of the proposed as well as of other 121 models found in the literature was performed, using 1-min resolution data. In this analysis, the elevation of the stations was not an essential factor when determining the parameters of the new model. The performance assessment was done using several statistical indicators and a Global Performance Index (GPI), which is a composite indicator that simplifies the analysis. The best performing model for the arid (AR) climate zone was found to be the model proposed by Abal et al. [67], while the best model for the high albedo (HA) climate zone is the model proposed by Mondol et al. [47]. The best model for the temperate (TM) and tropical (TR) climate zones is the model proposed in this work. It was also found that the best overall performing model (highest GPI in the more significant number of stations and highest skill score) was the model proposed in this work followed by the model by Mondol et al. [47]. This work helps to identify the best model that uses only the  $K_t$  according to the climate zone, thus allowing to easily estimate the diffuse horizontal irradiance (DHI) and consequently the direct normal irradiance (DNI), based only on global horizontal irradiance (GHI) measurements. The proposed model is also a tool for long-term data series quality control and gap-filling.

## Acknowledgements

Edgar F.M. Abreu acknowledges the support of the FCT (the Portuguese Science and Technology Foundation) through the grant with reference SFRH/BD/136433/2018. The authors acknowledge the funding provided by the European Union through the European Regional Development Fund, included in the COMPETE 2020 (Operational Program Competitiveness and Internationalization) through the ICT project (UID/GEO/04683/2013) with the reference POCI-01-0145-FEDER-007690 and the project DNI-Alentejo with reference ALT20-03-0145-FEDER-000011. The authors also acknowledge to the personnel who maintain the Baseline Surface Radiation Network radiometric stations.

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