

Determination of mixing-layer height by ground-based remote sensing

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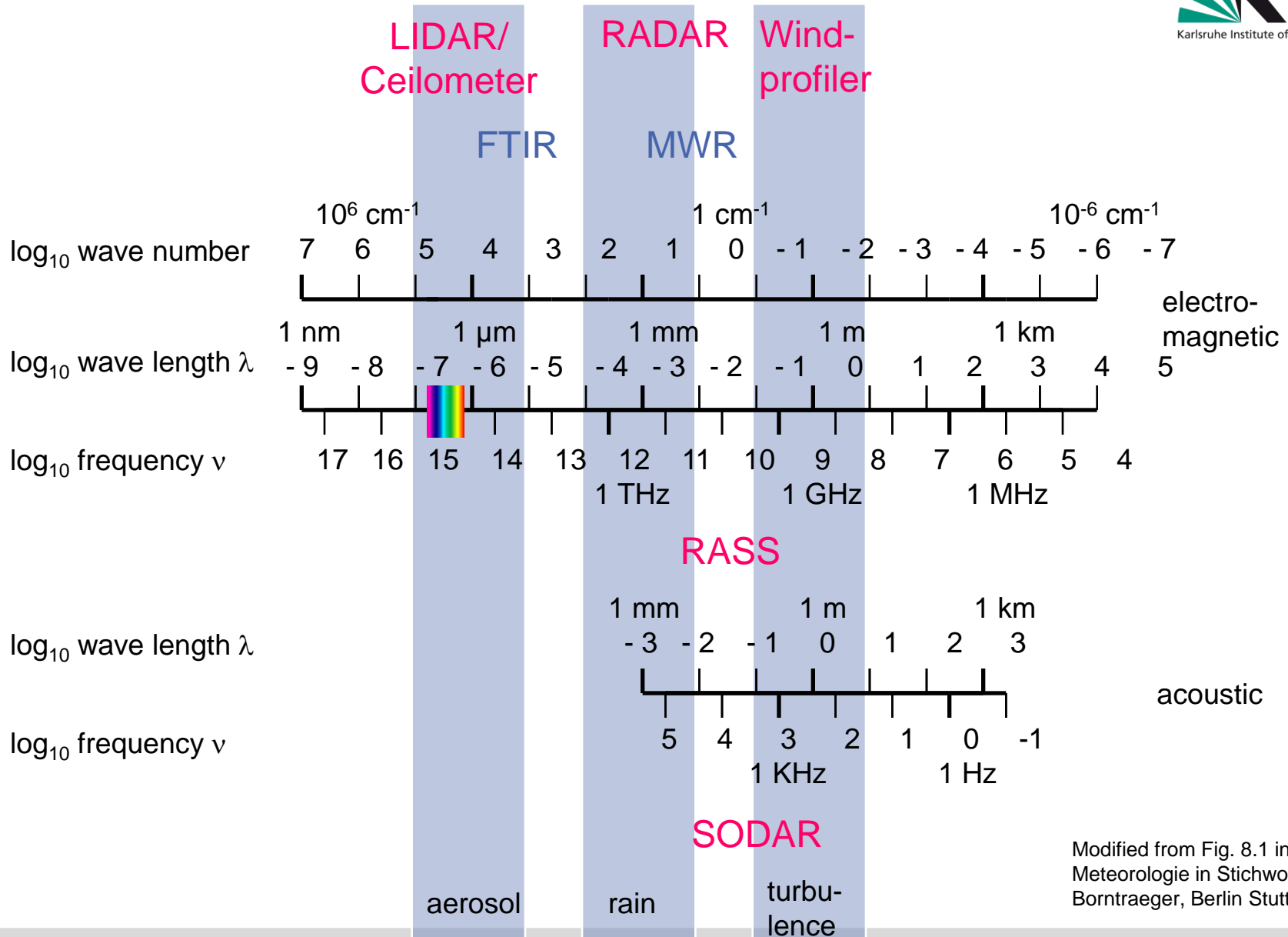
Basic remote sensing techniques

name	principle	spatial resolution	direction	type
RADAR	backscatter, electro-magnetic pulses, fixed wave length	profiling	scanning, slanted	active, monostatic
SODAR	backscatter, acoustic pulses, fixed wave length	profiling	fixed, slanted, vertical	active, usually monostatic
LIDAR	backscatter, optical pulses, fixed wave length(s)	profiling	scanning, fixed, horizontal, slanted, vertical	active, monostatic
RASS	backscatter, acoustic, electro-magnetic, fixed wave length	profiling	fixed, vertical	active, monostatic
	absorption, infrared, spectrum	path-averaging	fixed, horizontal, slanted	active, bistatic or passive
FTIR	emission, infrared, spectrum	path-averaging	fixed, horizontal, slanted	passive
DOAS	absorption, optical, fixed wave lengths	path-averaging	fixed, horizontal	active, bistatic
radiometry	electro-magnetic, fixed wave length(s)	averaging, profiling	fixed, scanning, slanted, vertical	passive
tomography	travel time, acoustic, fixed wave length	horizontal distribution	fixed, horizontal	active, multiple emitters and receivers

subject of this talk

subject of this talk

Typical frequency bands for remote sensing of the atmosphere



Modified from Fig. 8.1 in „
Meteorologie in Stichworten“,
Borntraeger, Berlin Stuttgart 2000

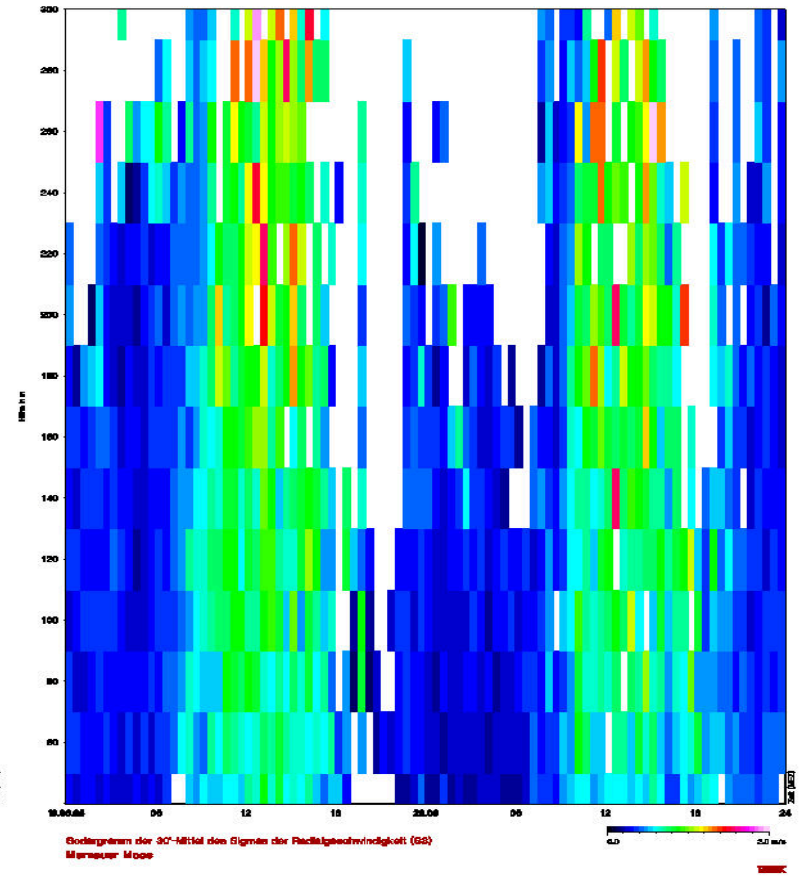
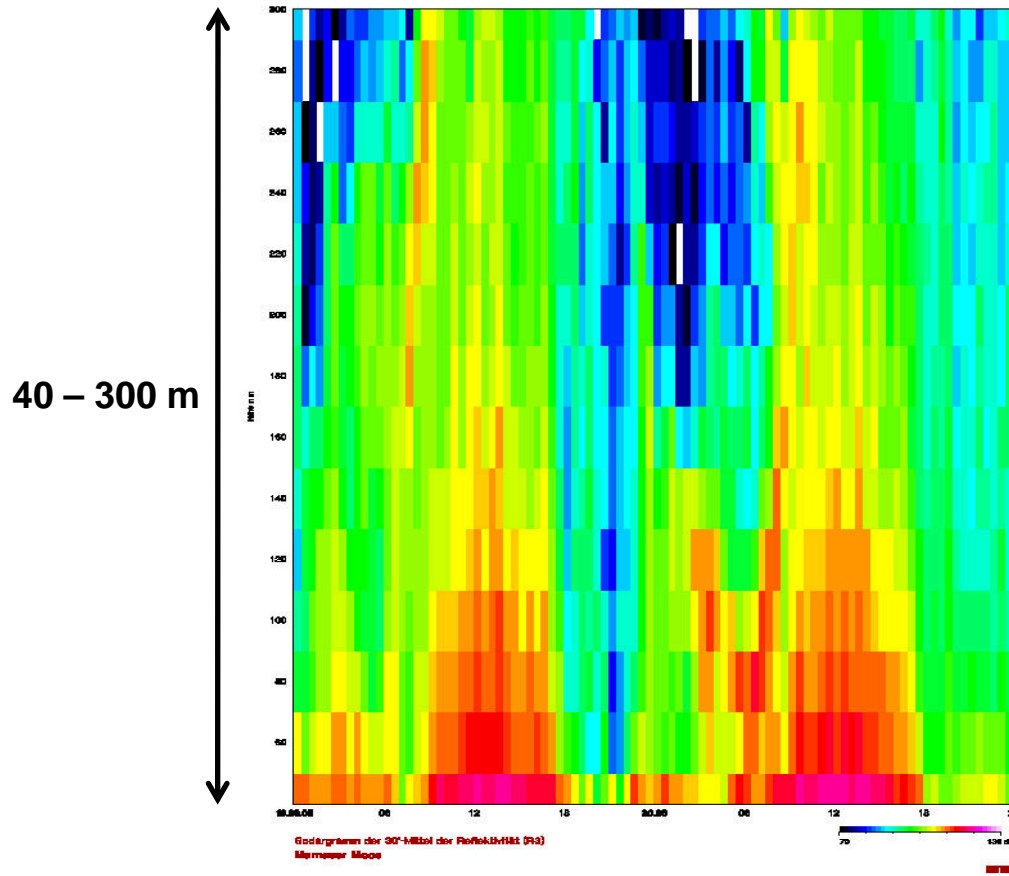
SODAR

algorithms for mixing-layer height

Sample plot SODAR (convective BL at daytime)

acoustic backscatter intensity

sigma w

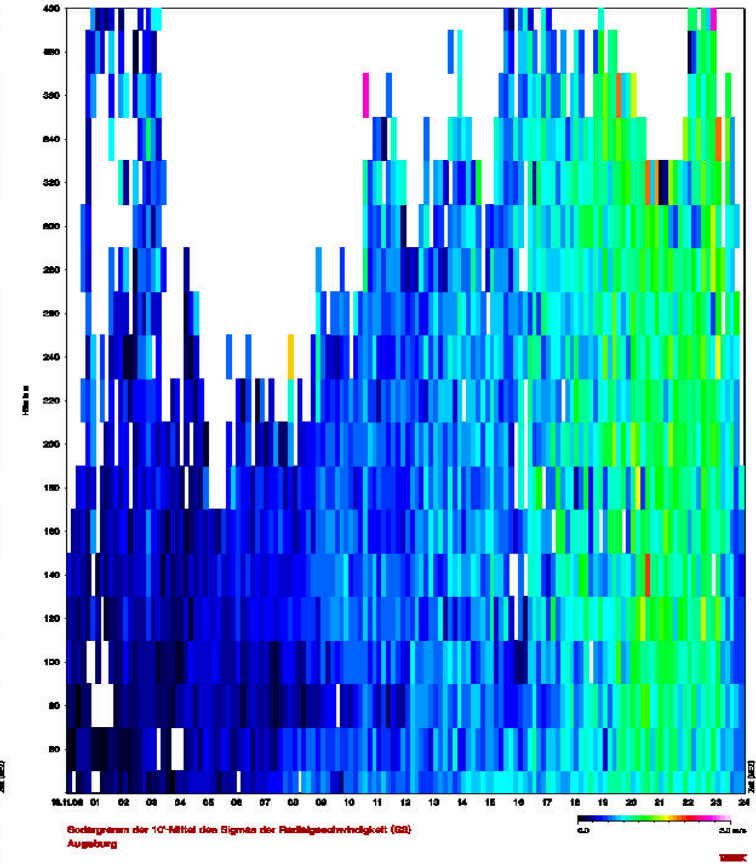
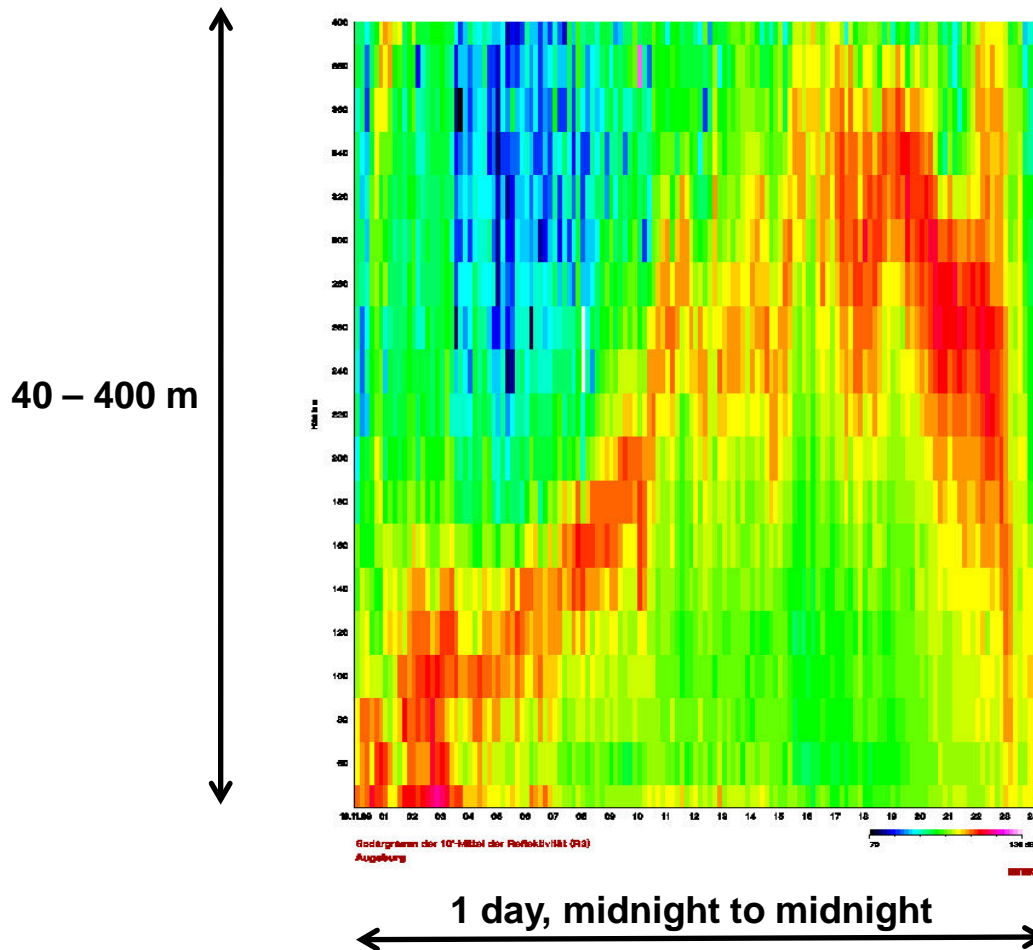


2 days, midnight to midnight

Sample plot SODAR (lifted inversion)

acoustic backscatter intensity

sigma w

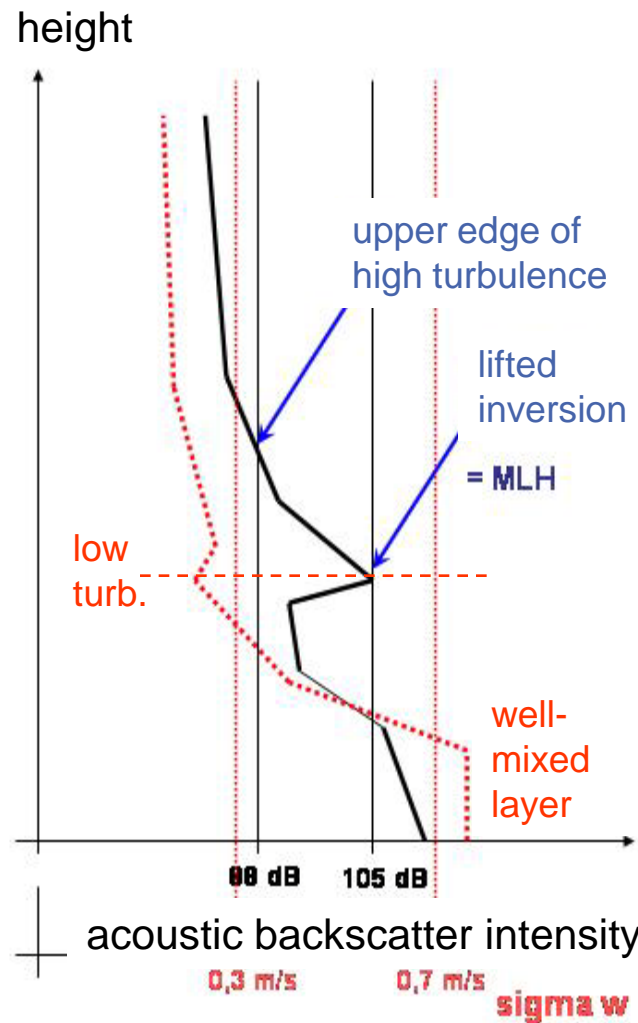


Algorithms to detect MLH from SODAR data

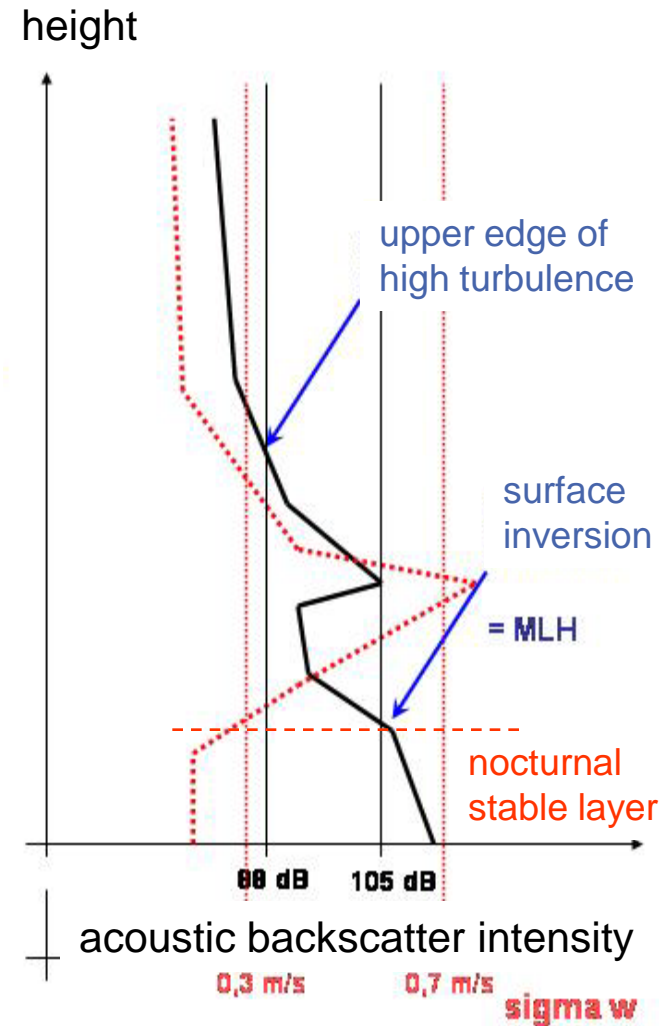
criterion 1:
 upper edge of high turbulence

criterion 2:
 surface and lifted inversions

$$MLH = \text{Min} (C1, C2)$$



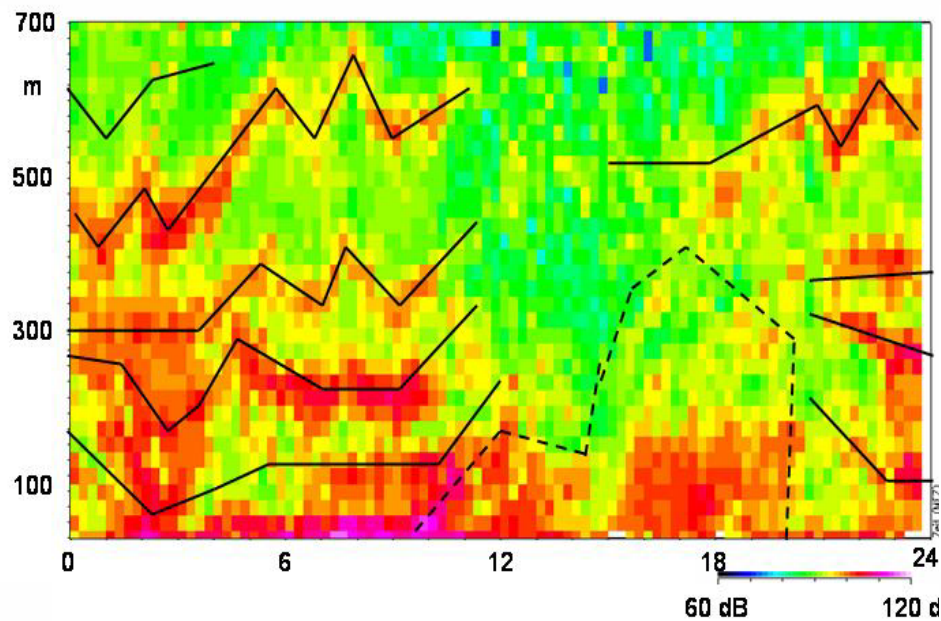
example 1: daytime



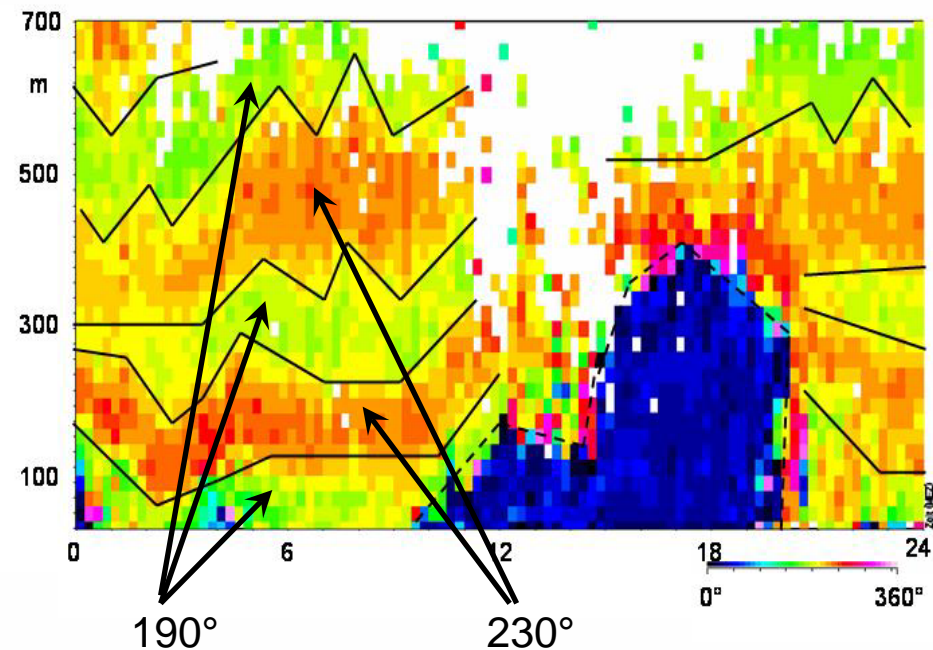
example 2: night-time

SODAR measurements in a wintry Alpine valley

29 January 2006



backscatter intensity



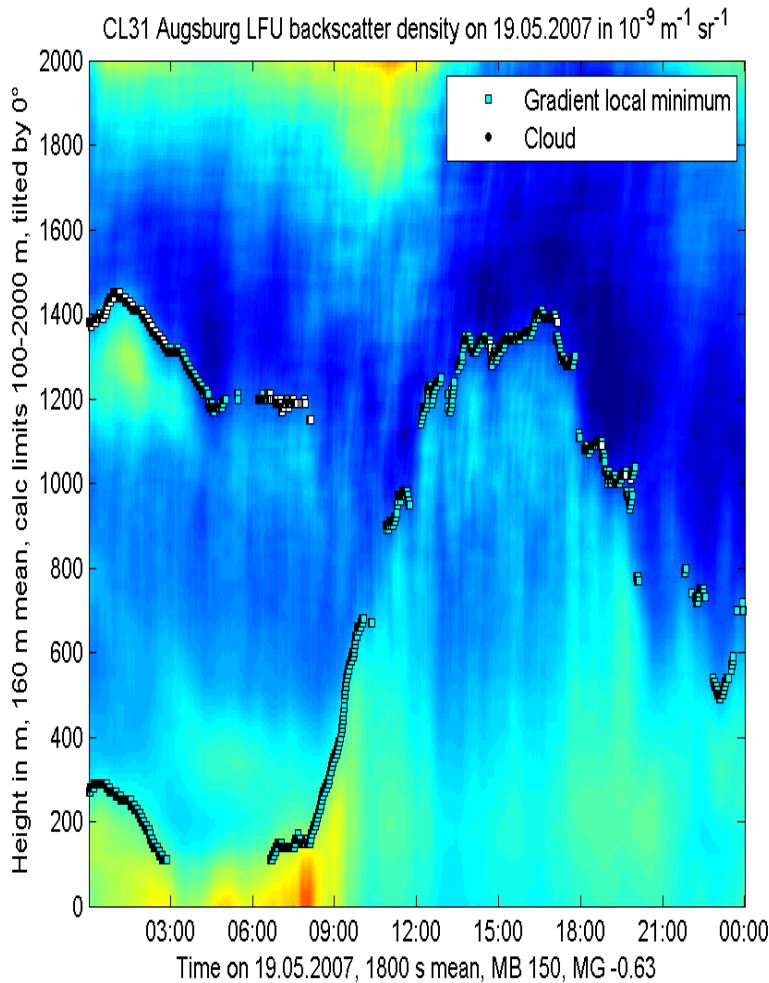
wind direction

Ceilometer

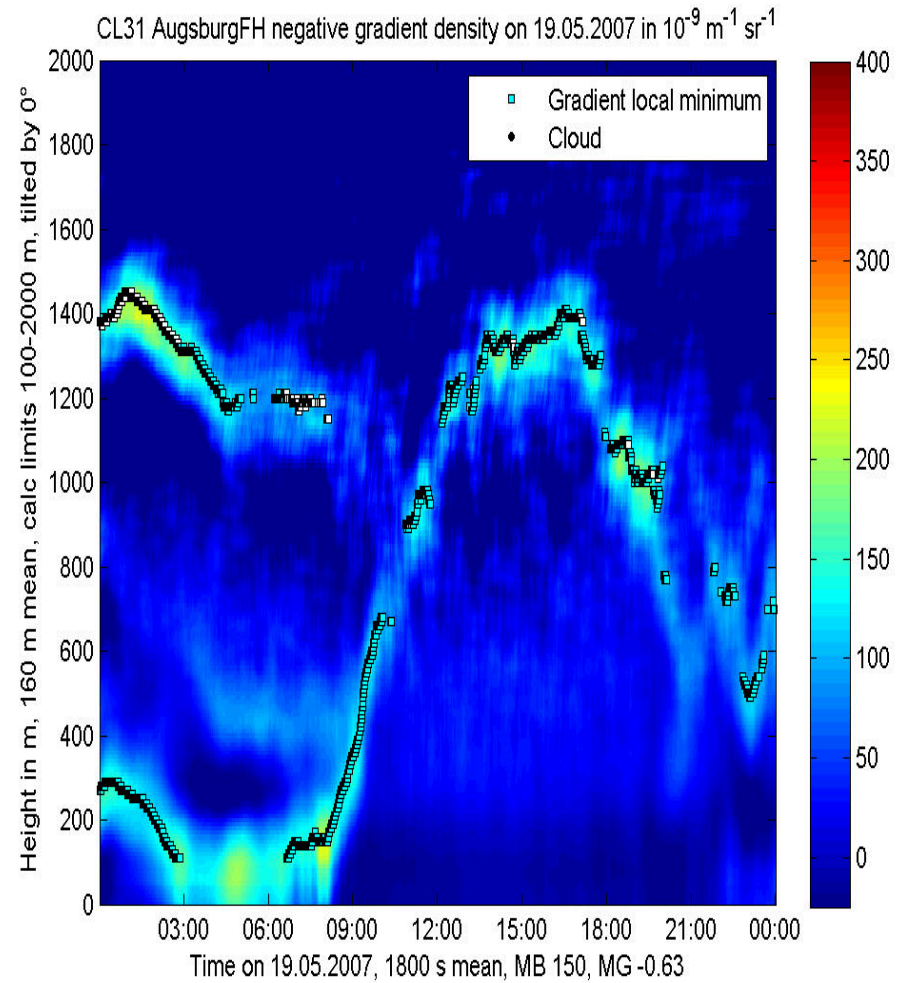
algorithms for mixing-layer height

Sample plot ceilometer (convective BL at daytime)

optical backscatter intensity



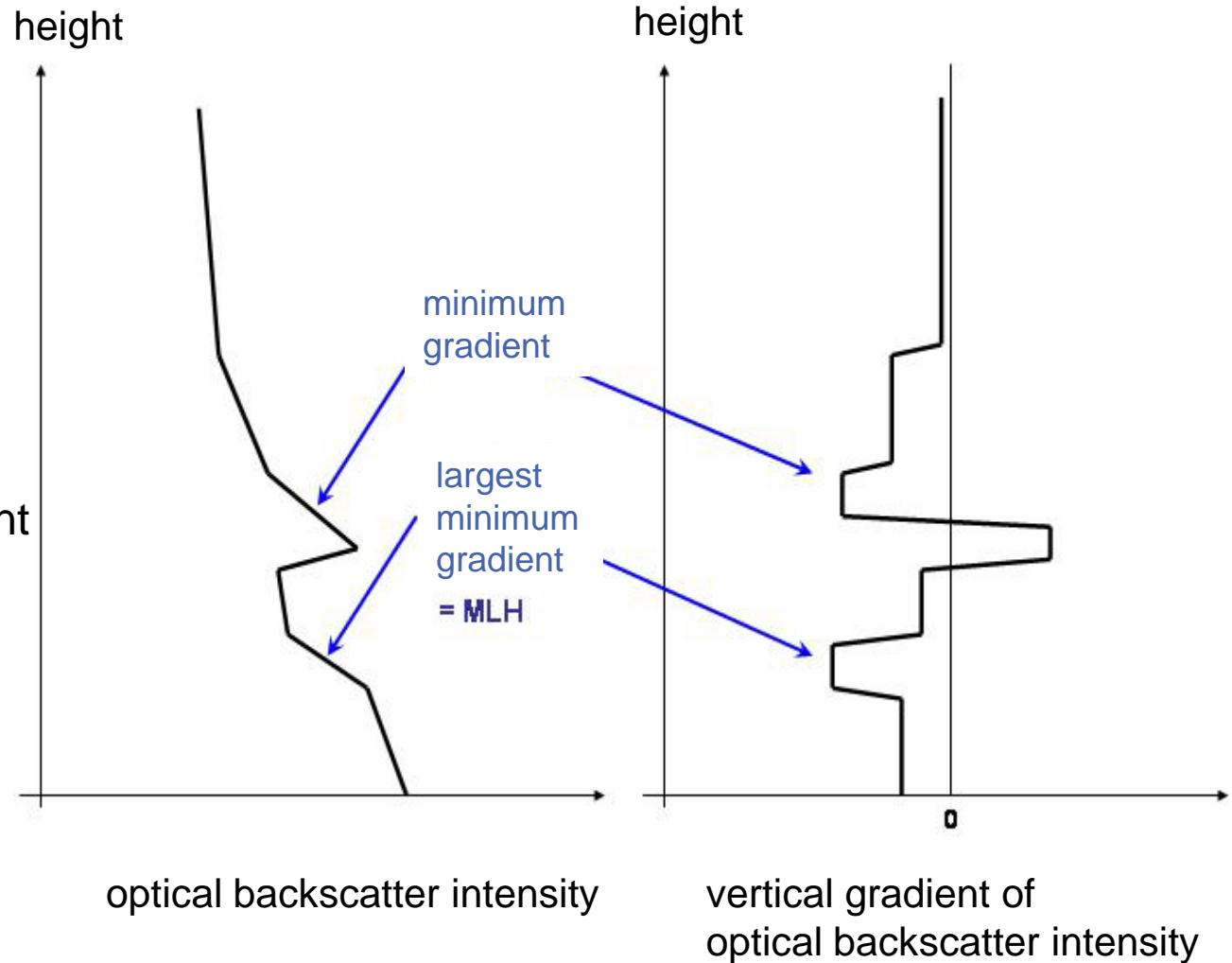
negative vertical gradient of optical backscatter intensity



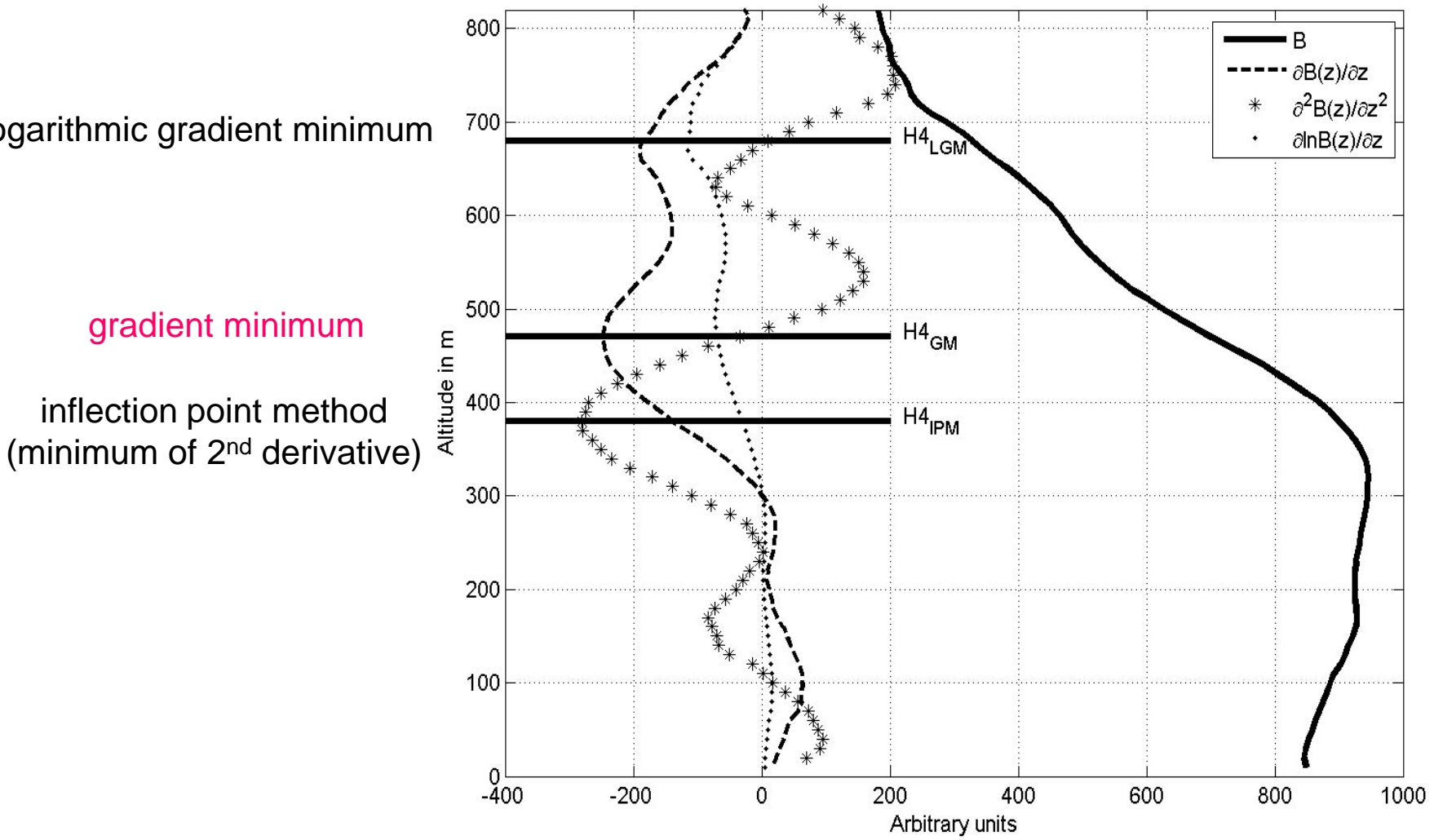
Algorithms to detect MLH from Ceilometer-Daten

criterion

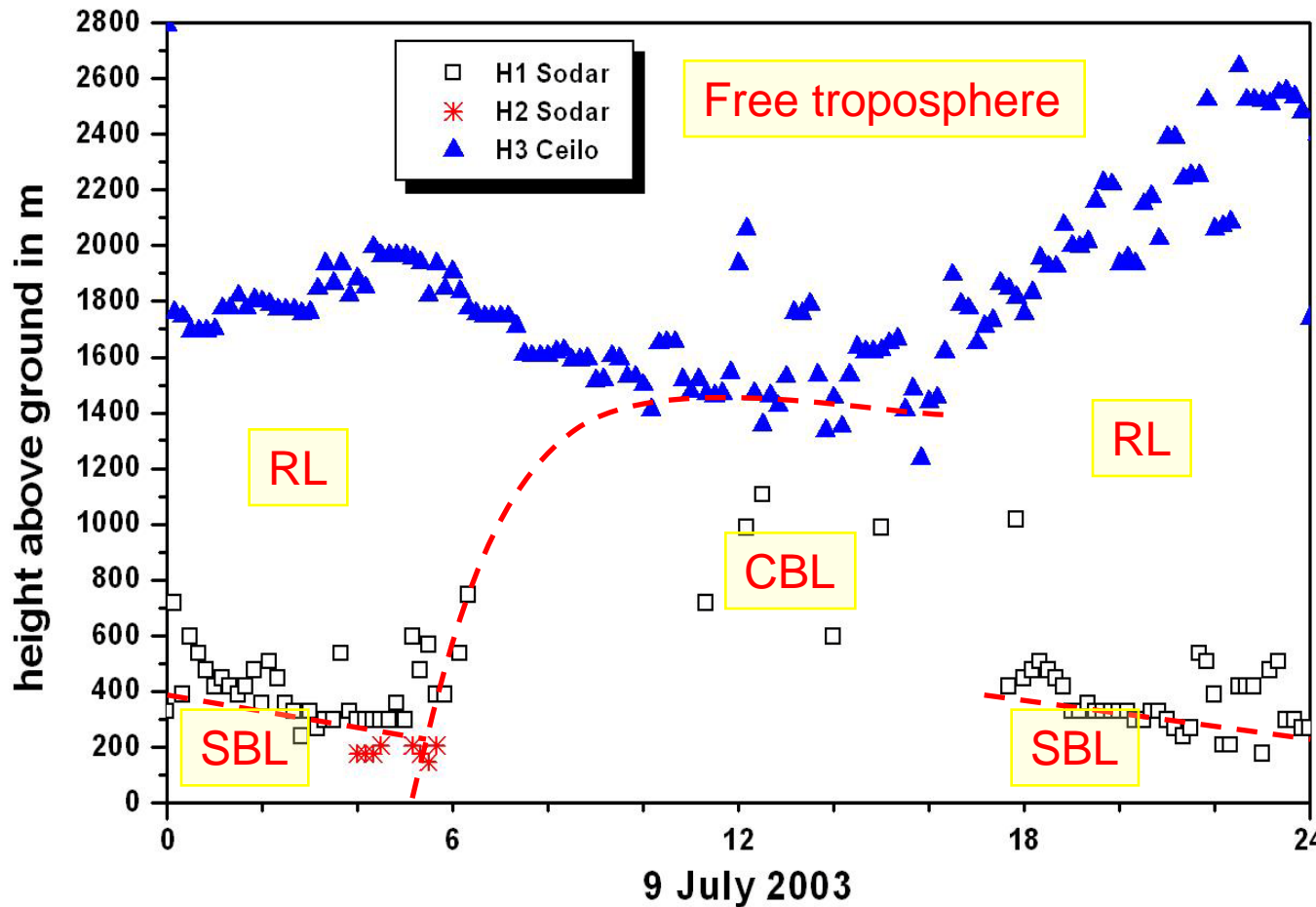
minimal vertical gradient of backscatter intensity (the most negative gradient)



Different gradient methods (see Sicard et al. 2006, BLM 119, 135-157)



Diurnal variation of mixing-layer height from SODAR and Ceilometer data (Budapest)



SBL:

stable boundary layer (usually at night and in winter)

CBL:

convective boundary layer (usually at daytime due to strong insolation)

RL:

residual layer (usually at night-time)



RASS

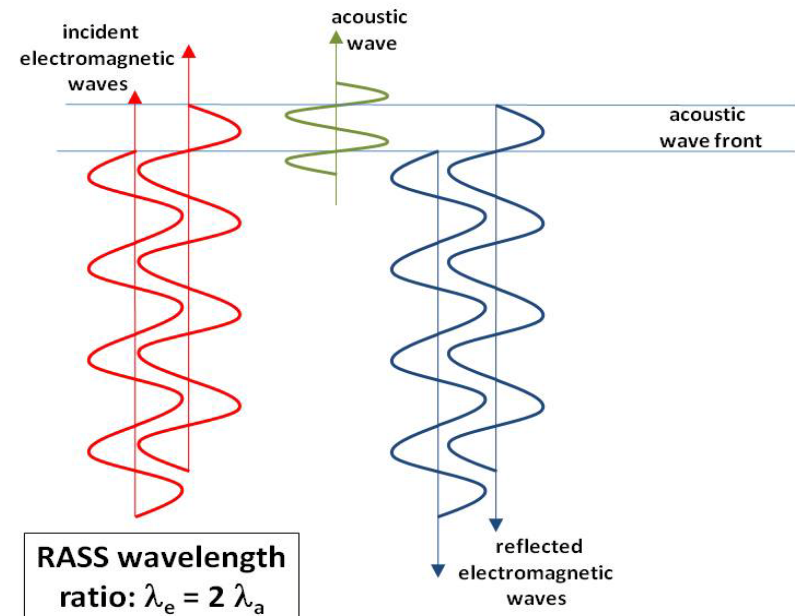
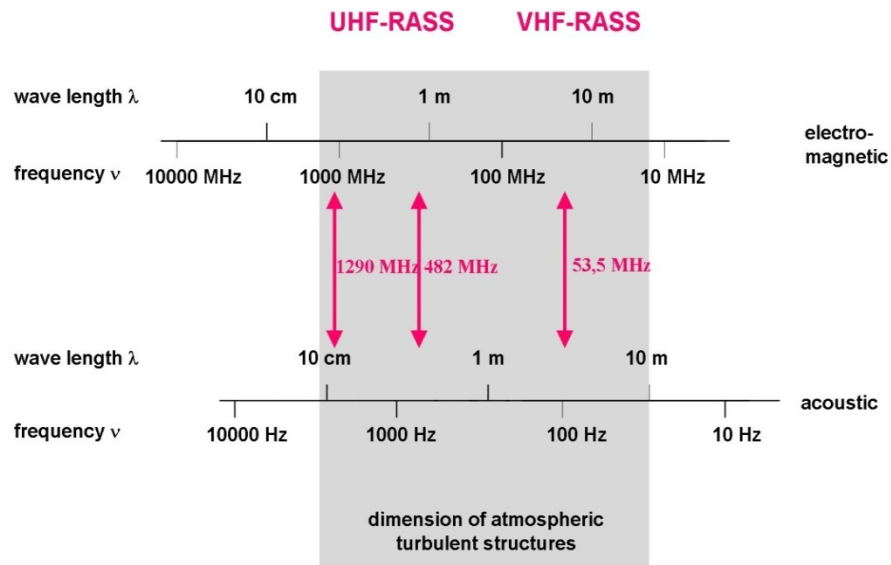
principles of operation

examples

RASS: frequencies

**Bragg condition:
acoustic wavelength = $\frac{1}{2}$ electro-magnetic wavelength**

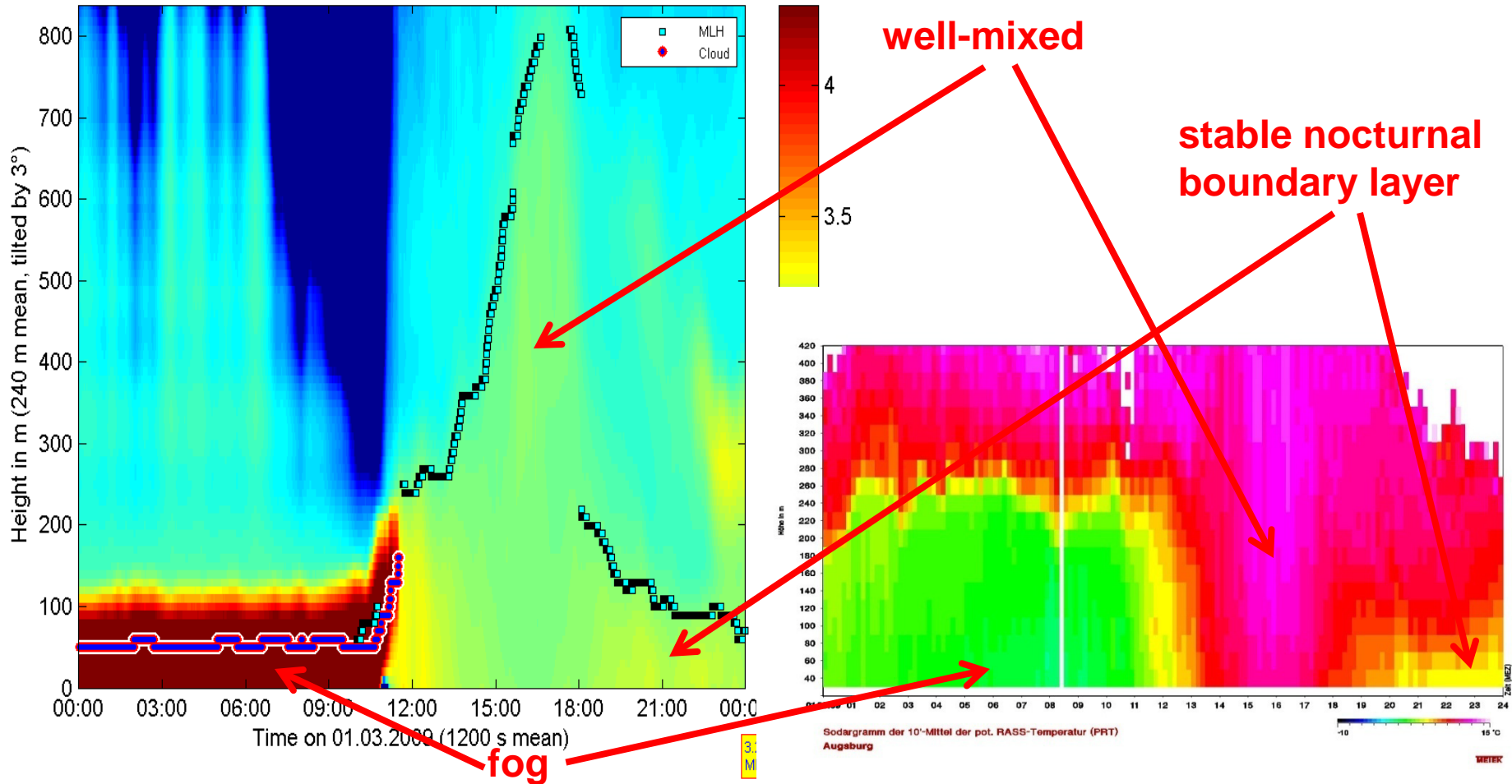
electro-magnetic - acoustic frequency pairs for RASS devices



temperature profile and pollution

comparison of RASS data (potential temperature, right)
with aerosol backscatter from a ceilometer (left)

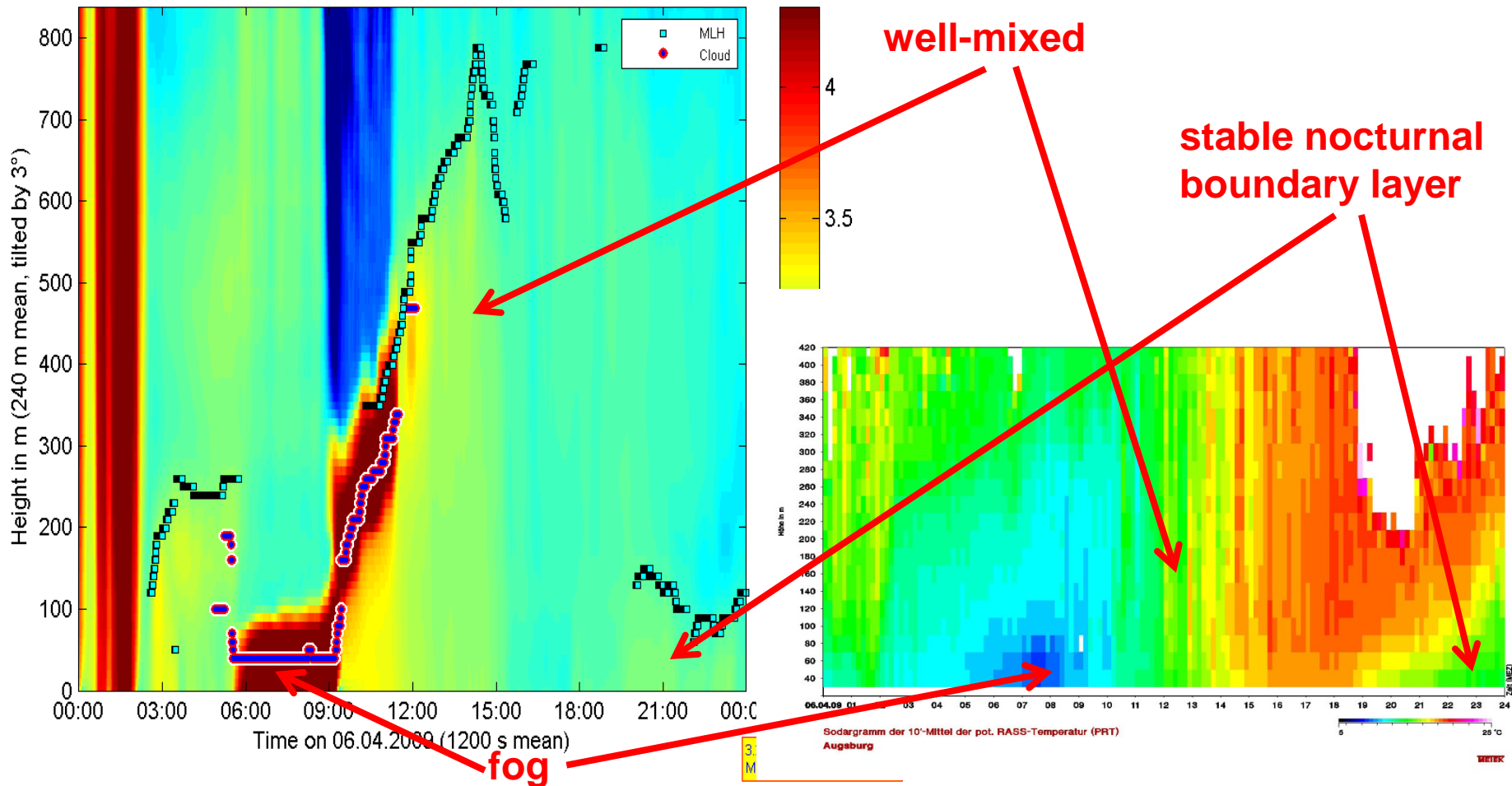
CL31 Augsburg AVA log₁₀ of backscatter with MLH on 01.03.2009 in 10⁻⁹ m⁻¹ sr⁻¹



temperature profile and pollution

comparison of RASS data (potential temperature, right)
with aerosol backscatter from a ceilometer (left)

CL31 Augsburg AVA log₁₀ of backscatter with MLH on 06.04.2009 in 10⁻⁹ m⁻¹ sr⁻¹



Overview on methods using ground-based remote sensing for the derivation of the mixing-layer height

method	short description
acoustic ARE method	analysis of acoustic received echo intensity profiles
“ HWS method	analysis of horizontal wind speed profiles
“ VWV method	analysis of vertical wind variance profiles
“ EARE method	analysis of acoustic backscatter intensity and vertical wind variance profiles (enhanced acoustic received echo method)
optical threshold method	detection of a given backscatter intensity threshold
“ gradient method	analysis of optical backscatter intensity profiles
“ idealised backscatter method	analysis of optical backscatter intensity profiles
“ wavelet method	analysis of optical backscatter intensity profiles
“ variance method	analysis of optical backscatter intensity profiles
acoustic / electro-magnetic	ARE method applied to sodar and wind profiler data
acoustic / optical	EARE method plus gradient method
electro-magnetic / electro-magnetic	combination of a sodar-RASS and a wind profiler RASS: analysis of the vertical temperature profile plus analysis of the electro-magnetic backscatter intensity profile
acoustic / in situ	ARE method plus in-situ surface flux measurement
RASS	analysis of the temperature profile from the measured speed of sound

Conclusions:

RASS directly delivers temperature profiles. MLH, inversions, and stable layers can easily be detected, wind profiles are additionally available. **Only remote system that measures inversion strengths.** Does not work properly with high wind speeds.

SODAR detects temperature fluctuations and gradients, but no absolute temperature. Inversions and stable layers can indirectly be inferred with a MLH algorithm. Does not work properly with perfectly neutral stratification, with very high wind speeds, and during stronger precipitation events.

Ceilometer detects aerosol distribution and water droplets. It has to be assumed that the aerosol follows the thermal structure of the atmosphere. Inversions and MLH can indirectly be inferred with a MLH algorithm. Does not work properly in extreme clear (aerosol-free) air and during precipitation events and fog.

SODAR:

Asimakopoulos, D.N., C.G. Helmis, J. Michopoulos, 2004: Evaluation of SODAR methods for the determination of the atmospheric boundary layer mixing height. - Meteor. Atmos. Phys. 85, 85–92.



Beyrich, F., 1997: Mixing height estimation from sodar data – a critical discussion. - Atmos. Environ. 31, 3941–3953.

Ceilometer:

Schäfer, K., S.M. Emeis, A. Rauch, C. Münkel, S. Vogt, 2004: Determination of mixing-layer heights from ceilometer data. In: Remote Sensing of Clouds and the Atmosphere IX. Schäfer, K., A. Comeron, M. Carleer, R.H. Picard, N. Sifakis (Eds.), Proc. SPIE, Bellingham, WA, USA, Vol. 5571, 248–259.

Sicard, M., C. Pérez, F. Rocadenbosch, J.M. Baldasano, D. García-Vizcaino, 2006: Mixed-Layer Depth Determination in the Barcelona Coastal Area From Regular Lidar Measurements: Methods, Results and Limitations. - Bound.-Lay. Meteor. 119, 135–157.

RASS:

Engelbart, D.A.M., J. Bange, 2002: Determination of boundary-layer parameters using wind profiler/RASS and sodar/RASS in the frame of the LITFASS project. Theor. Appl. Climatol. 73, 53–65.

Emeis, S., K. Schäfer, C. Münkel, 2009: Observation of the structure of the urban boundary layer with different ceilometers and validation by RASS data. Meteorol. Z., 18, 149-154. (Open access, freely available from <http://dx.doi.org/10.1127/0941-2948/2009/0365>)

Reviews:

Emeis, S., K. Schäfer, C. Münkel, 2008: Surface-based remote sensing of the mixing-layer height – a review. - Meteorol. Z., 17, 621-630. (Open access, freely available from <http://dx.doi.org/10.1127/0941-2948/2008/0312>)

Emeis, S., M. Harris, R.M. Banta, 2007: Boundary-layer anemometry by optical remote sensing for wind energy applications. - Meteorol. Z., 16, 337-347.

