

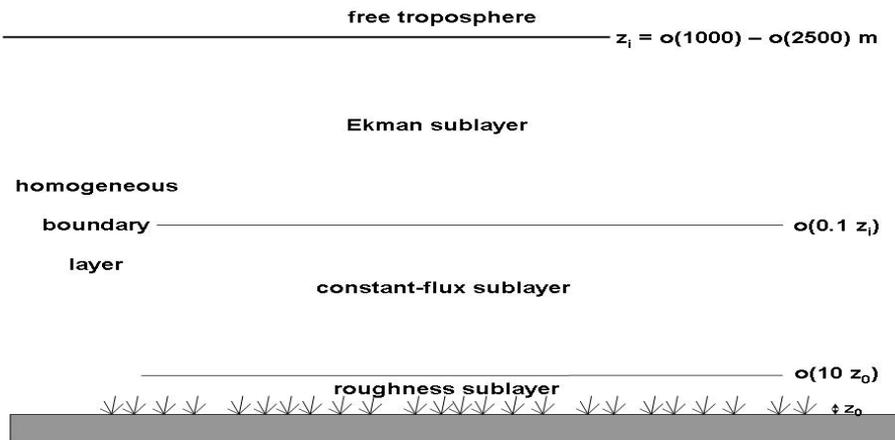
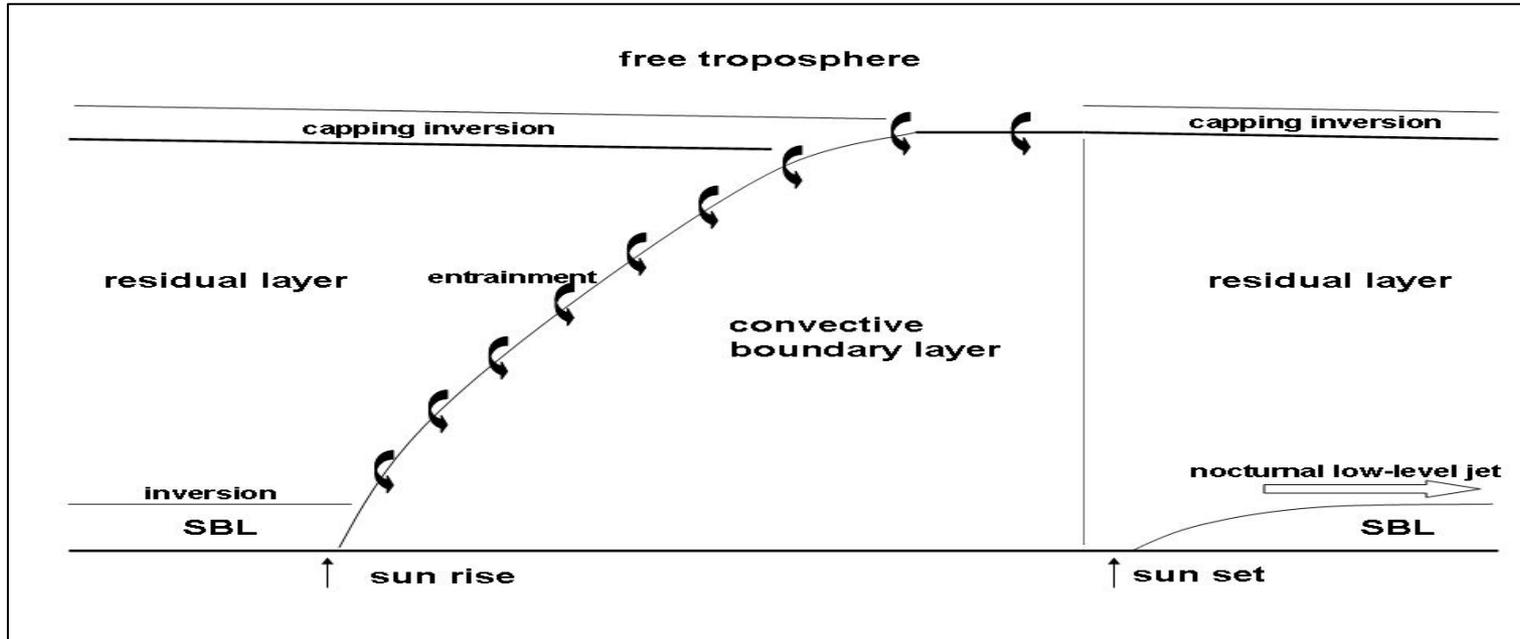
RASS-Sondierungen zur Verifikation von MLH-Bestimmungen aus SODAR- und Ceilometerdaten

Stefan Emeis

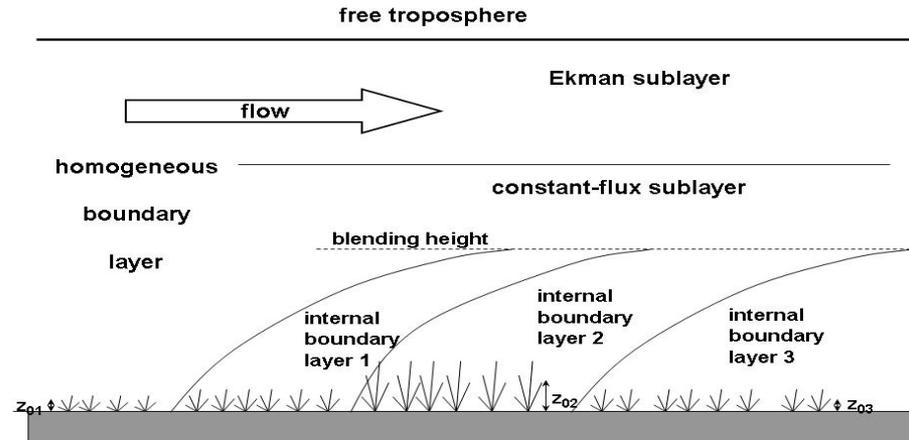
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zusammen mit: Klaus Schäfer, Christoph Münkler (Väisälä) und Peter Suppan

diurnal variation of PBL

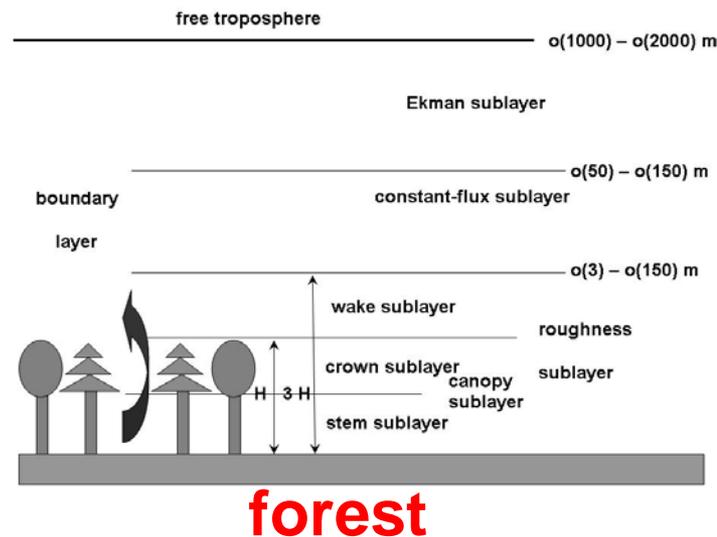
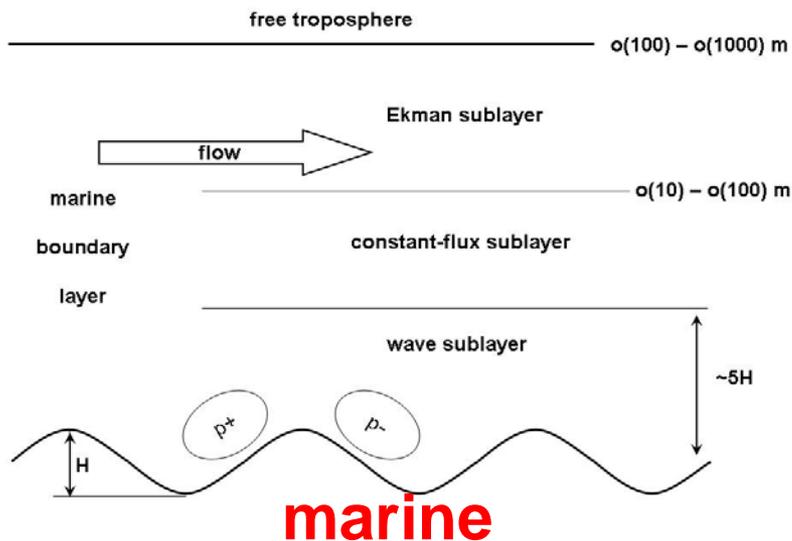
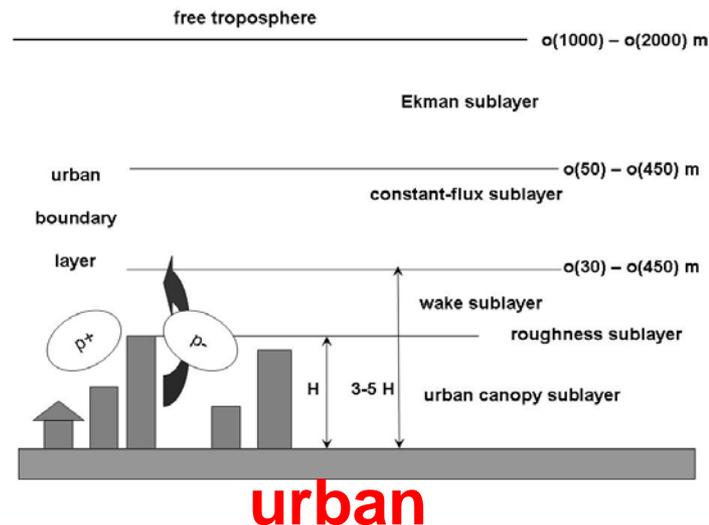
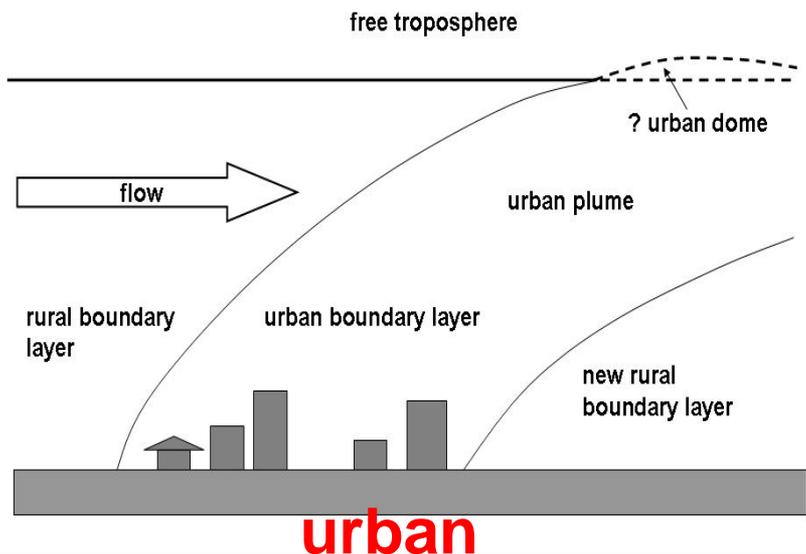


vertical structure of PBL



internal layers in PBL

special types of PBL



The vertical wind profile

logarithmic law

(with stability correction) $u(z) = (u_*/\kappa) (\ln(z/z_0) - \psi(z/L_*))$

power law

$$u(z) = u(z_A) (z/z_A)^n$$

New proposal
(Gryning et al. 2007)

$$u(z) = \frac{u_{*0}}{\kappa} \left(\ln \left(\frac{z}{z_0} \right) + \frac{z}{L_{MBL,N}} - \frac{z}{z_i} \left(\frac{z}{2L_{MBL,N}} \right) \right)$$

needs information on the PBL or mixing-layer height

Gryning et al., 2007: On the extension of the wind profile over homogeneous terrain beyond the surface boundary layer. Bound.-Lay. Meteorol., **124**, 251–268.

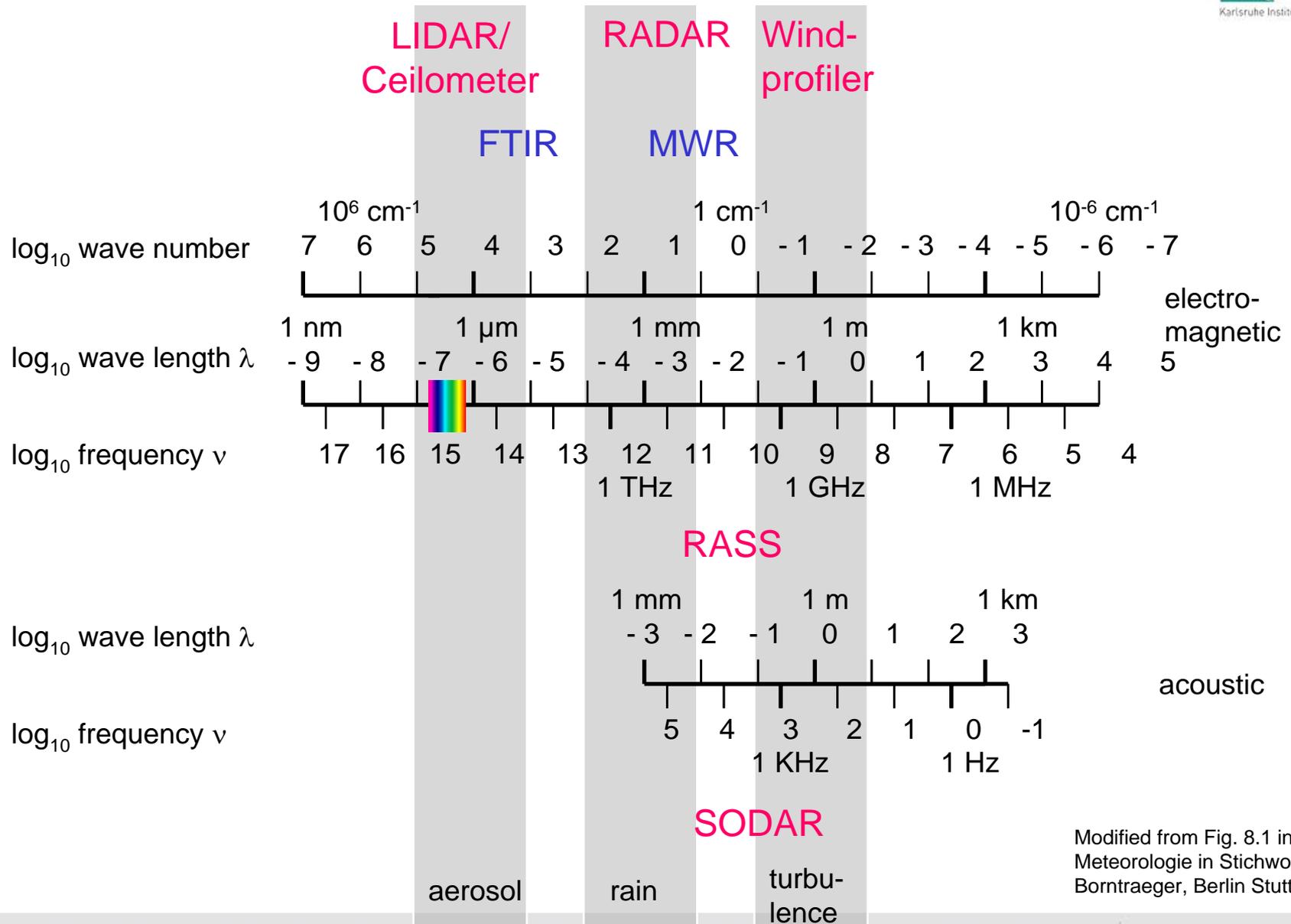
Basic remote sensing techniques

name	principle	spatial resolution	direction	type
RADAR	backscatter, electro-magnetic pulses, fixed wave length	profiling	scanning, slanted	active, monostatic
SODAR	backscatter, acoustic pulses, fixed wave length	profiling	fixed, slanted, vertical	active, usually monostatic
LIDAR	backscatter, optical pulses, fixed wave length(s)	profiling	scanning, fixed, horizontal, slanted, vertical	active, monostatic
RASS	backscatter, acoustic, electro-magnetic, fixed wave length	profiling	fixed, vertical	active, monostatic
	absorption, infrared, spectrum	path-averaging	fixed, horizontal, slanted	active, bistatic or passive
FTIR	emission, infrared, spectrum	path-averaging	fixed, horizontal, slanted	passive
DOAS	absorption, optical, fixed wave lengths	path-averaging	fixed, horizontal	active, bistatic
radiometry	electro-magnetic, fixed wave length(s)	averaging, profiling	fixed, scanning, slanted, vertical	passive
tomography	travel time, acoustic, fixed wave length	horizontal distribution	fixed, horizontal	active, multiple emitters and receivers

subject of this talk

subject of this talk

Typical frequency bands for remote sensing of the atmosphere



Modified from Fig. 8.1 in „
Meteorologie in Stichworten“,
Borntraeger, Berlin Stuttgart 2000

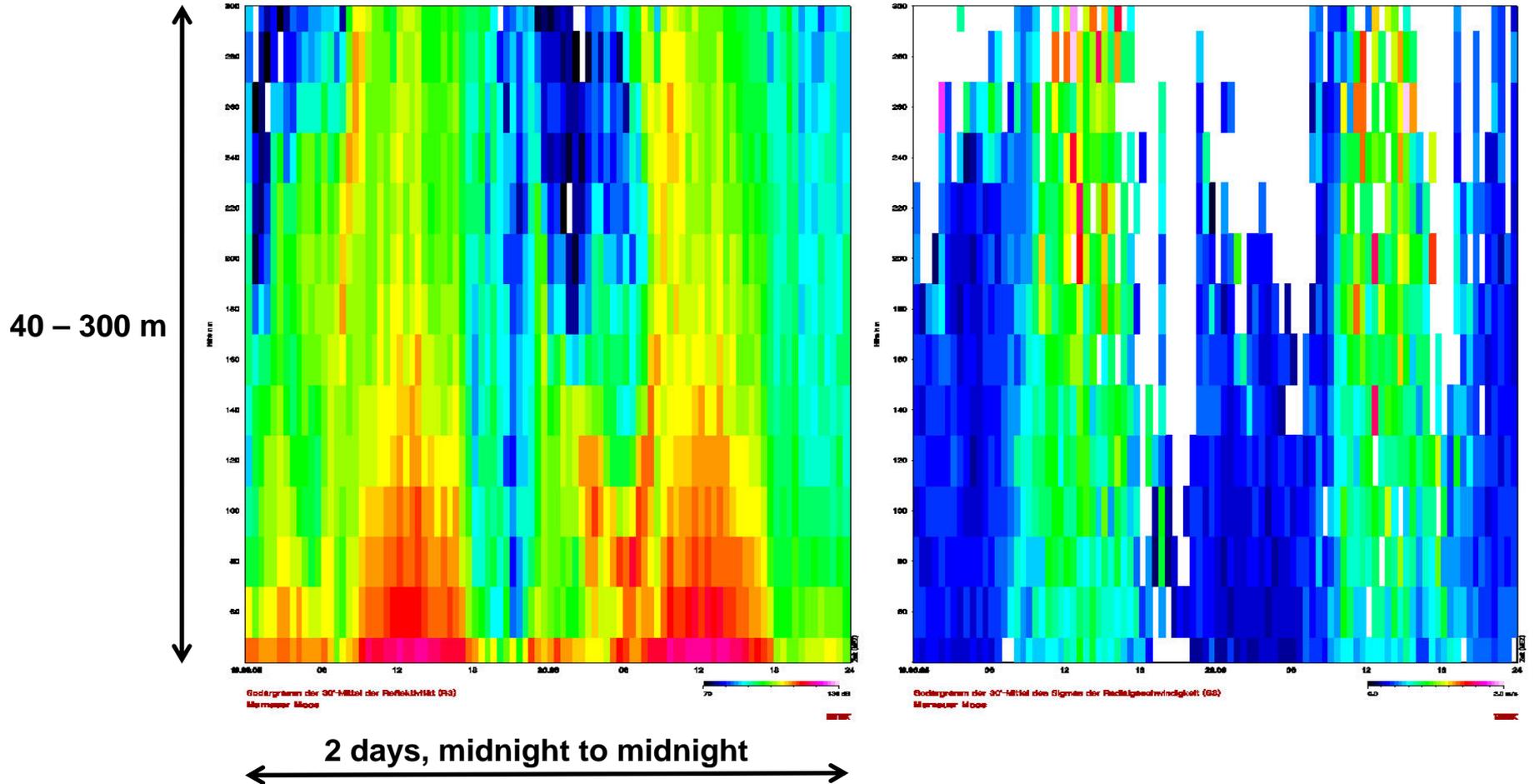
SODAR

algorithms for mixing-layer height

Sample plot SODAR (convective BL at daytime)

acoustic backscatter intensity

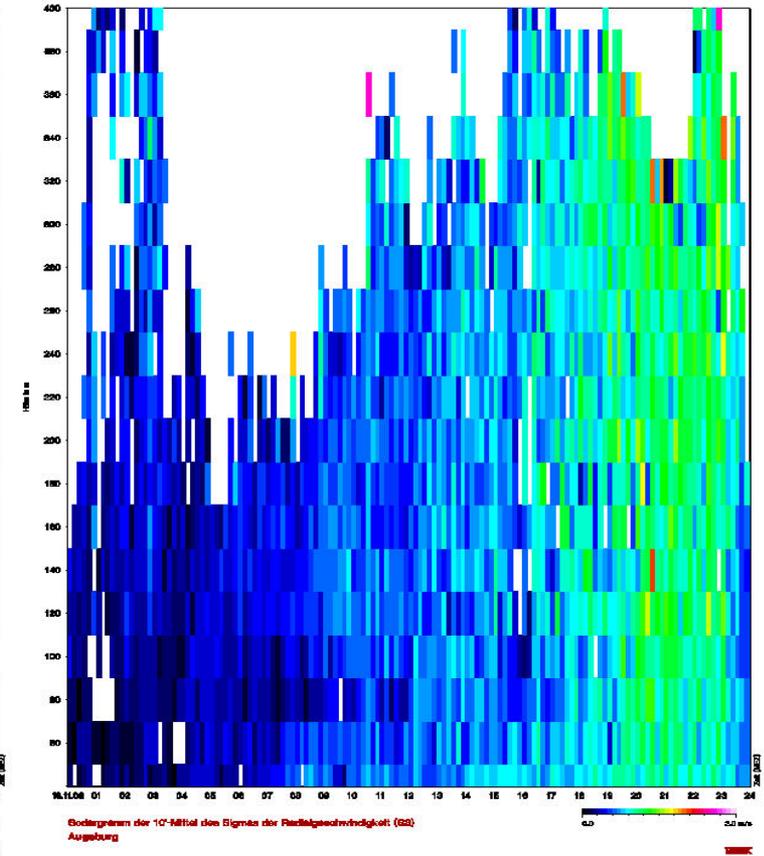
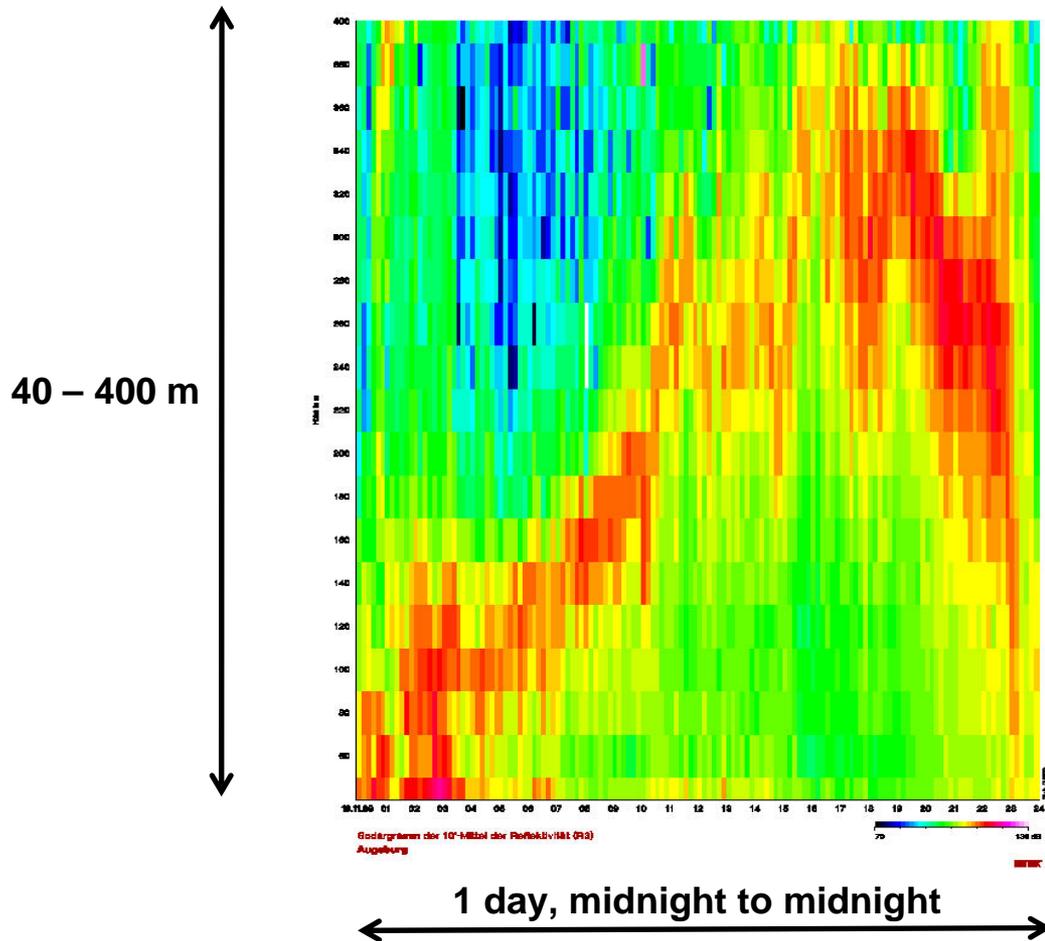
sigma w



Sample plot SODAR (lifted inversion)

acoustic backscatter intensity

sigma w

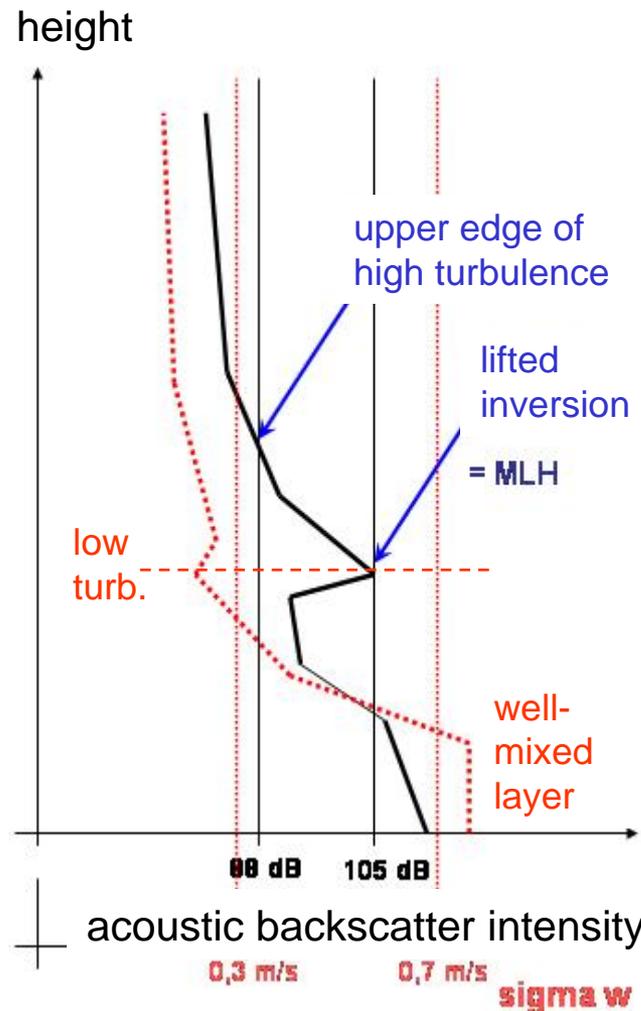


Algorithms to detect MLH from SODAR data

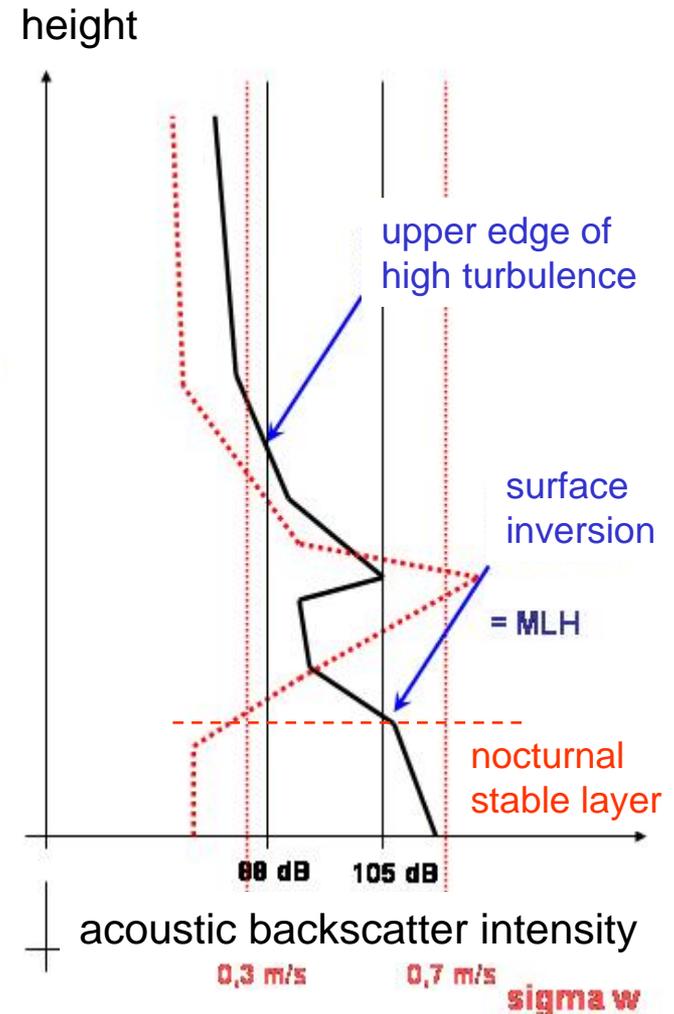
criterion 1:
 upper edge of high turbulence

criterion 2:
 surface and lifted inversions

MLH = Min (C1, C2)



example 1: daytime



example 2: night-time

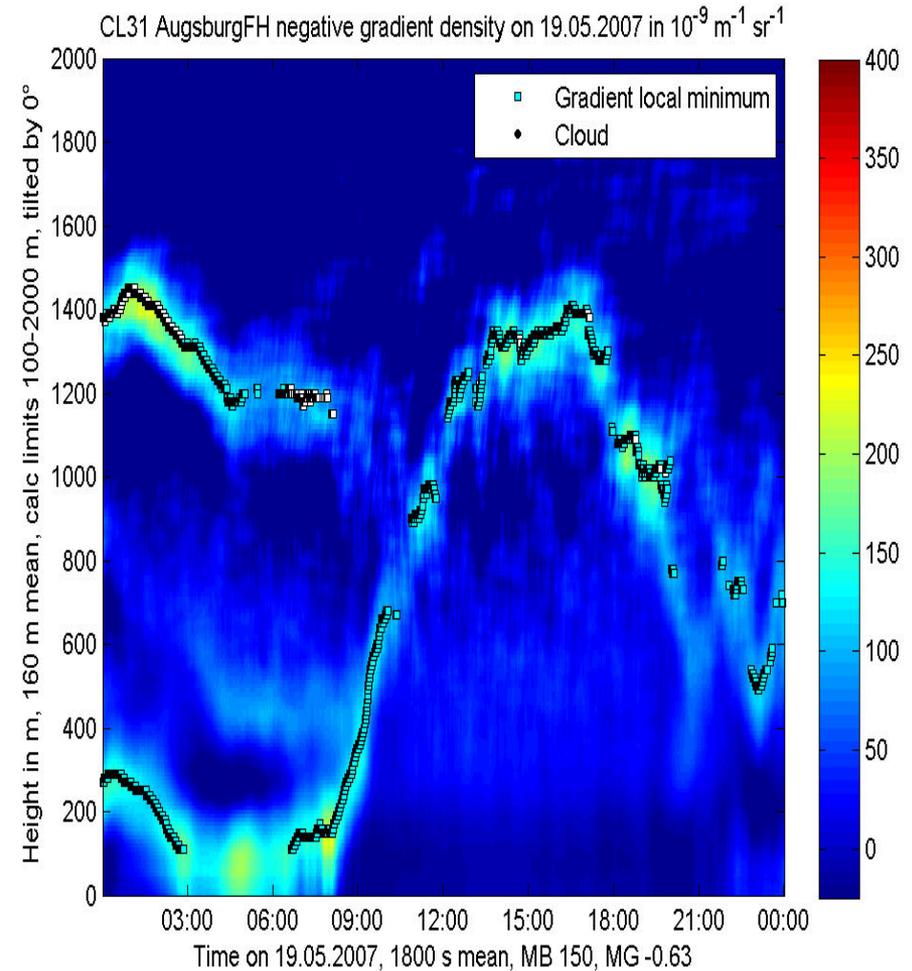
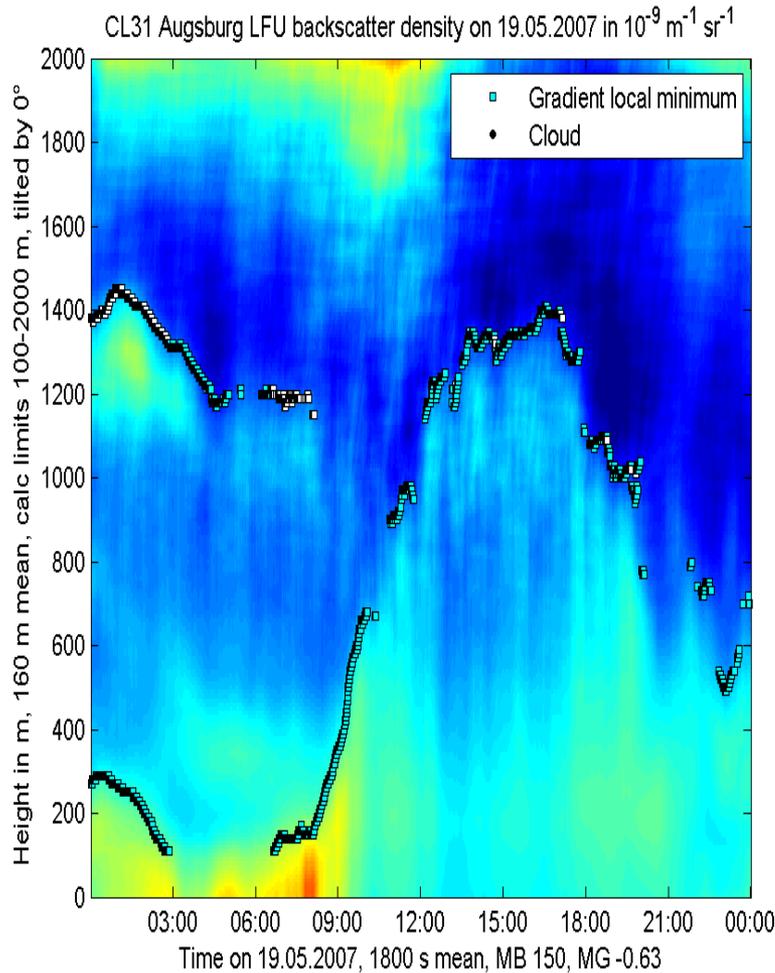
Ceilometer

algorithms for mixing-layer height

Sample plot ceilometer (convective BL at daytime)

optical backscatter intensity

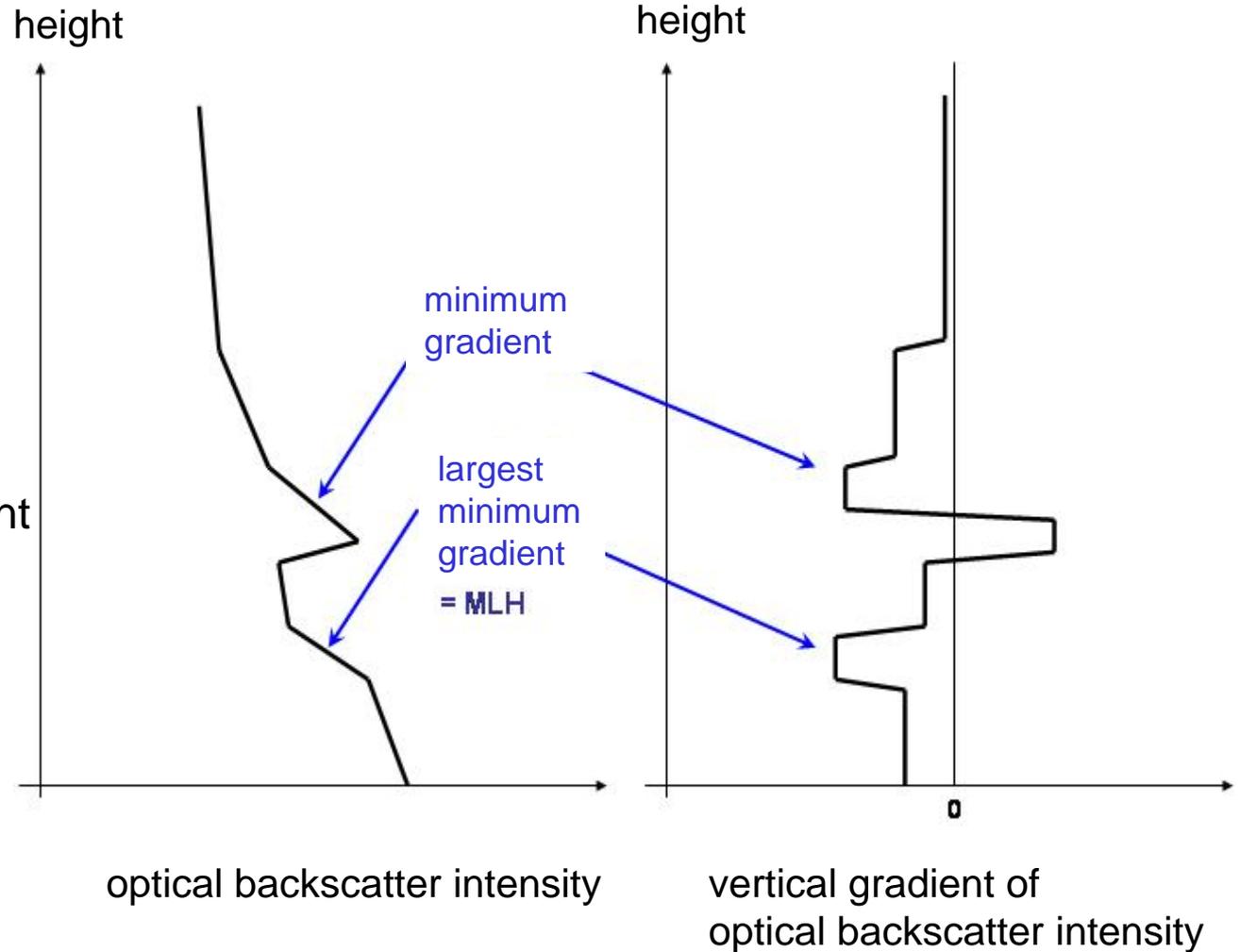
negative vertical gradient of optical backscatter intensity



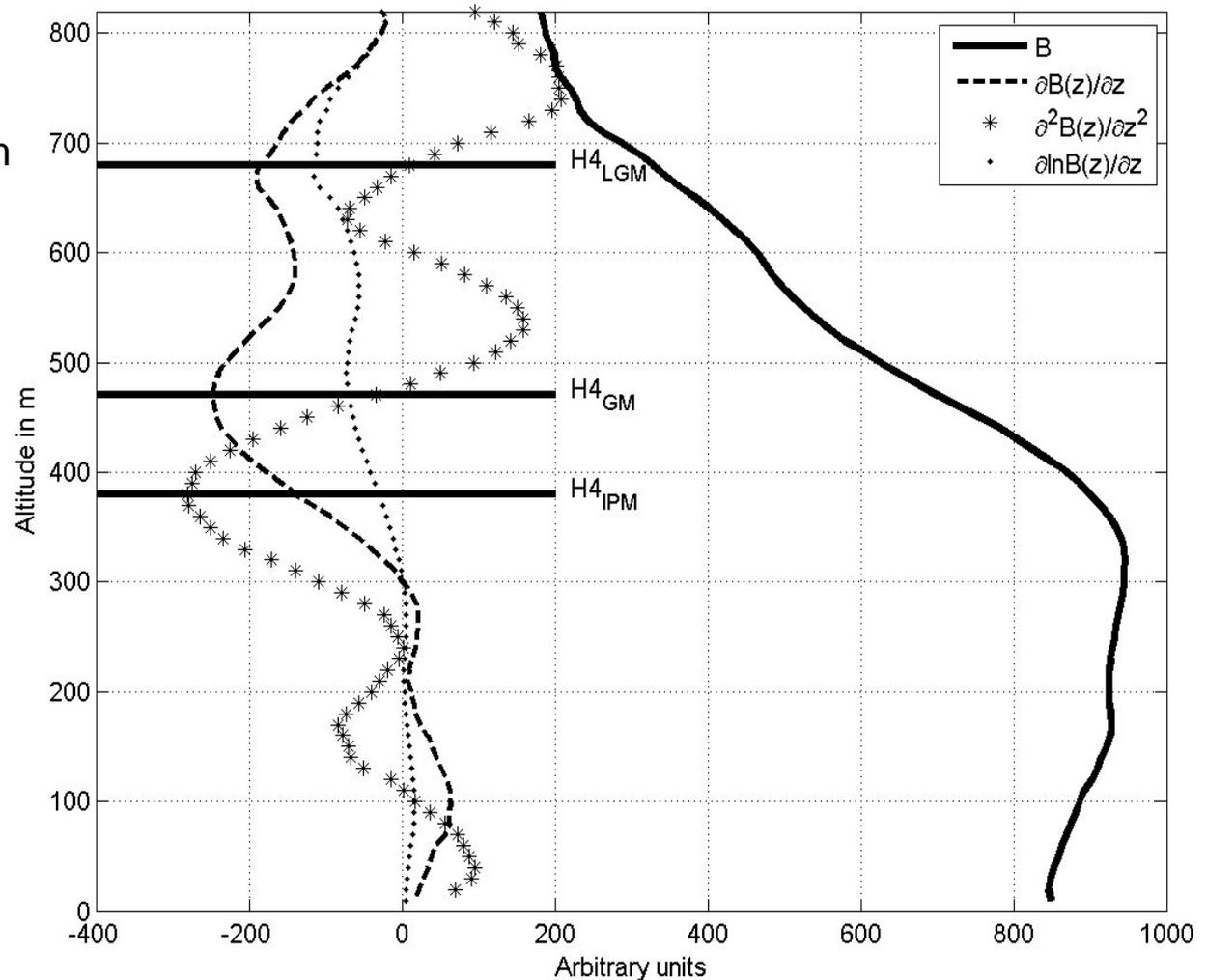
Algorithms to detect MLH from Ceilometer-Daten

criterion

minimal vertical gradient of backscatter intensity (the most negative gradient)



Different gradient methods (see Sicard et al. 2006, BLM 119, 135-157)



logarithmic gradient minimum

gradient minimum

inflection point method
(minimum of 2nd derivative)

RASS

Messprinzip

Vergleiche mit Ceilometer und SODAR

RASS (radio-acoustic remote sensing)

misst vertikale Temperaturprofile

Bragg-RASS: windprofiler plus akustische Komponente

Doppler-RASS: SODAR plus elektro-magnetische Komponente

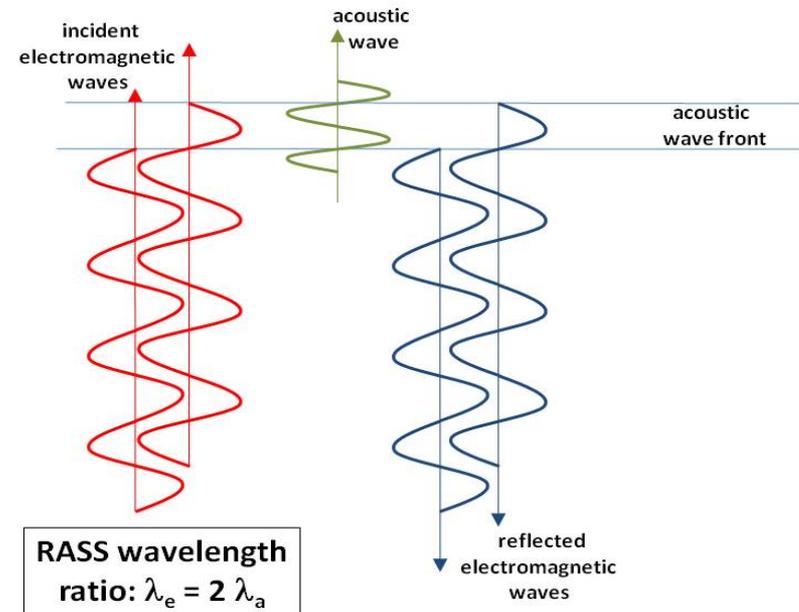
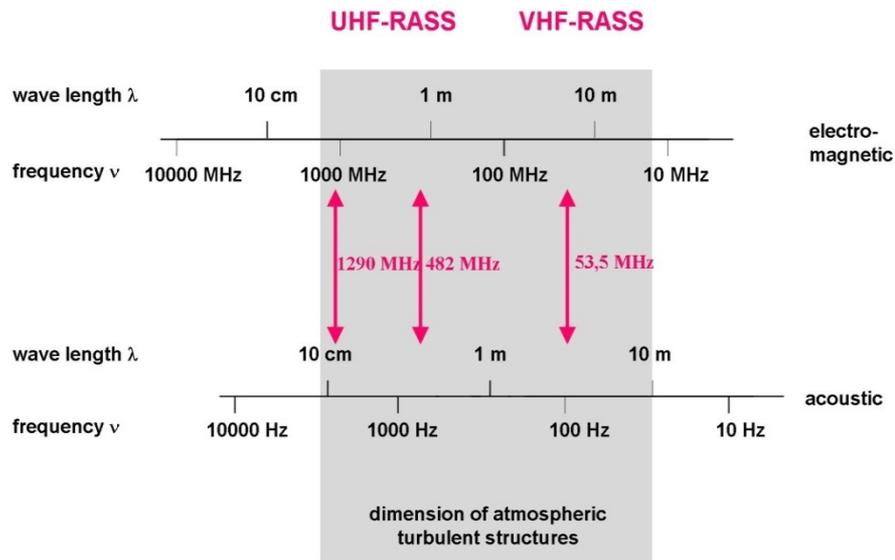
UHF RASS (Grenzschicht)

VHF RASS (Troposphäre)

RASS: Frequenzen

Bragg Bedingung: akustische Wellenlänge = $\frac{1}{2}$ elektro-magnetische Wellenlänge

electro-magnetic - acoustic frequency pairs for RASS devices





SODAR-RASS (Doppler-RASS)

(METEK)

akust. Frequ.: 1500 – 2200 Hz

Radio Frequ.: 474 MHz

Auflösung: 20 m

unterstes

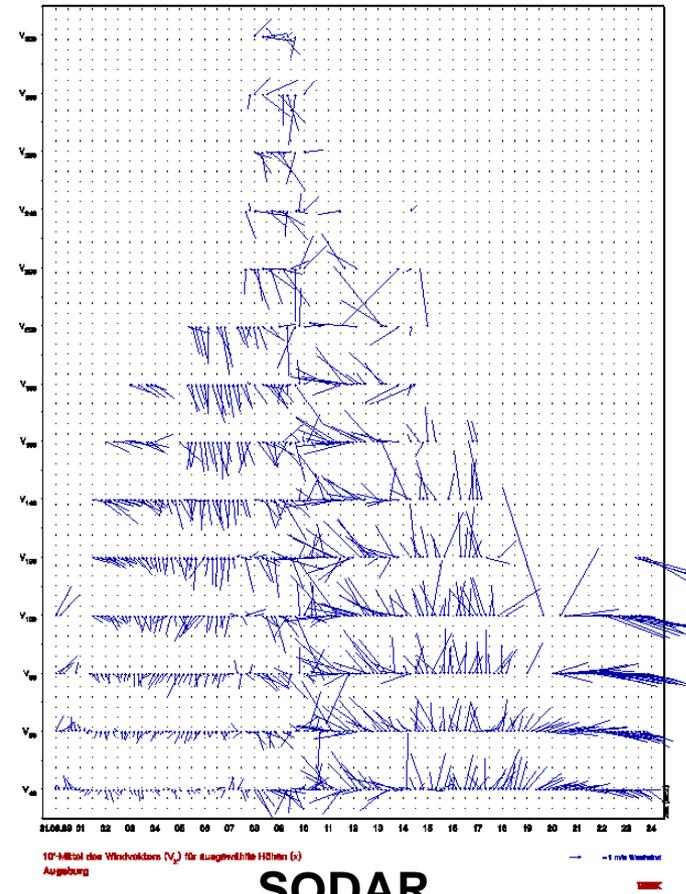
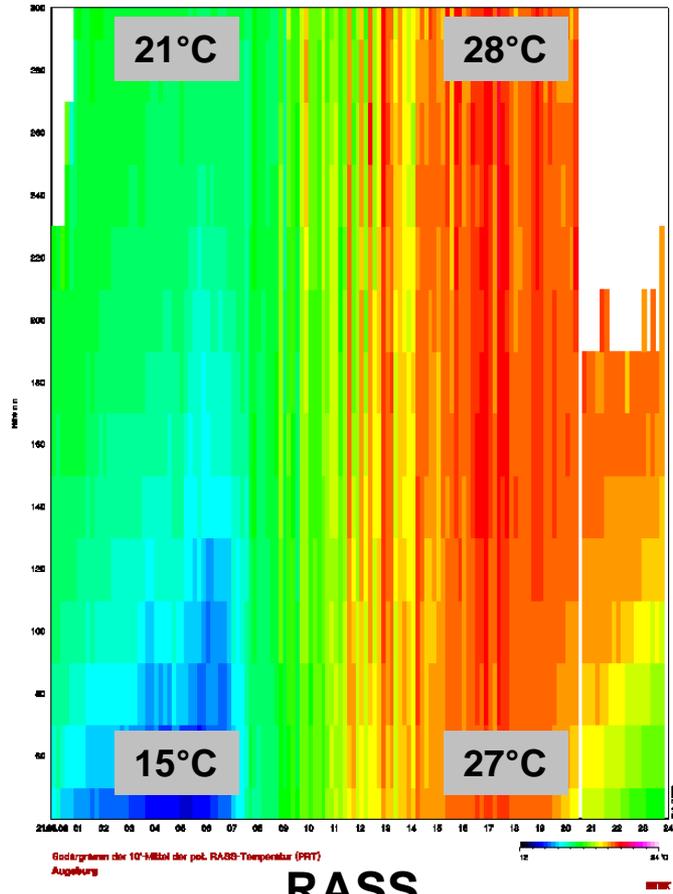
Messniveau: ca. 40 m

Reichweite: 540 m

Vergleich mit SODAR-Wind

Sommertag
potenzielle Temperatur (links), horizontaler Wind (rechts)

300 m

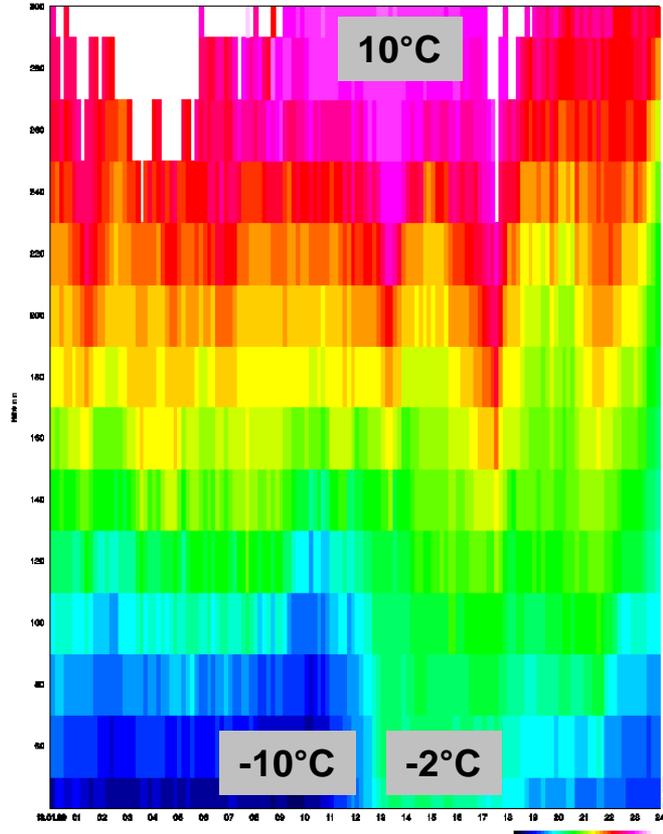


40 m

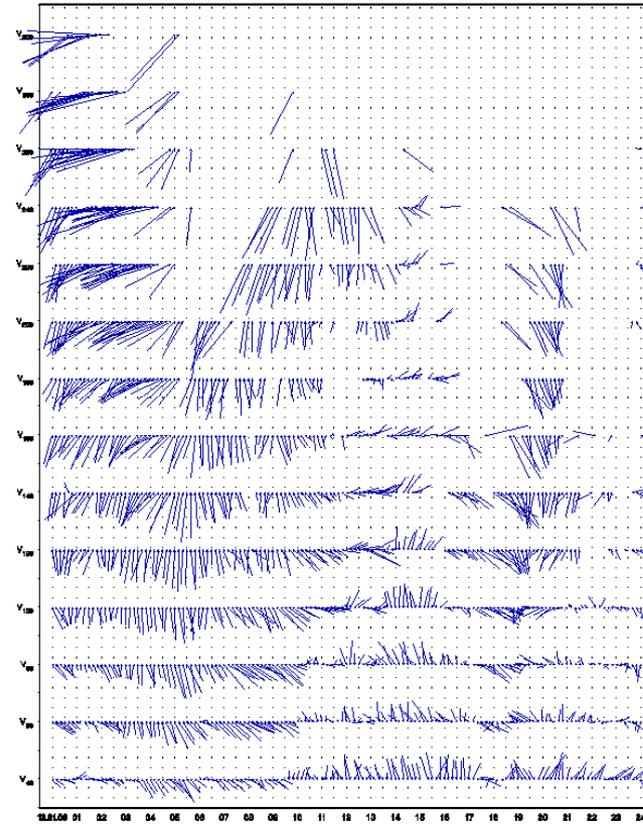
Vergleich mit SODAR-Wind

Wintertag
potenzielle Temperatur (links), horizontaler Wind (rechts)

300 m



RASS

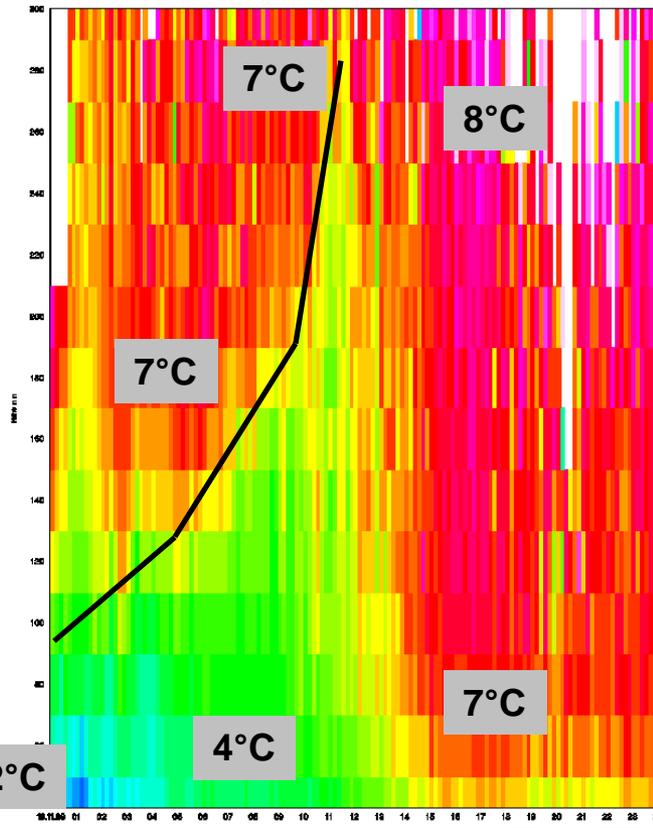


SODAR

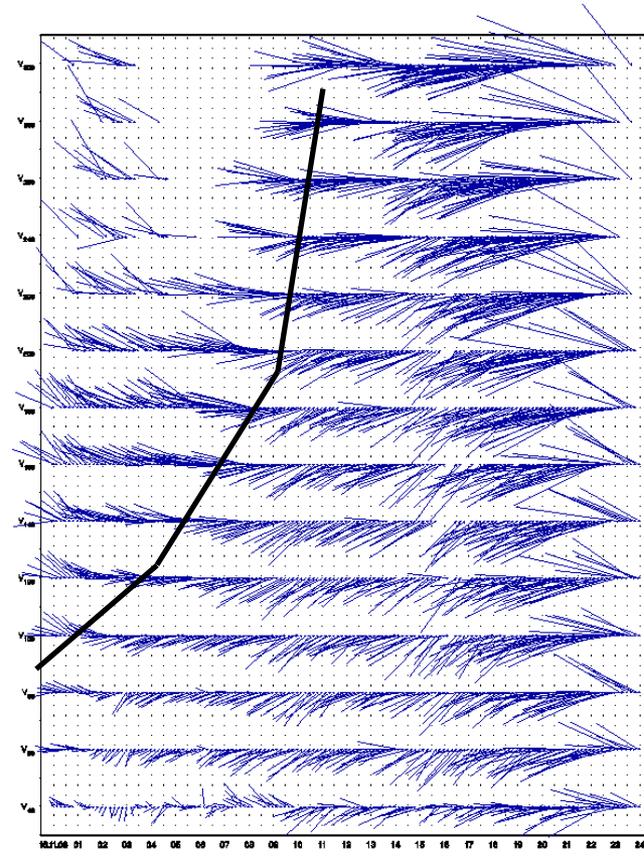
Vergleich mit SODAR-Wind

Inversion
potenzielle Temperatur (links), horizontaler Wind (rechts)

300 m

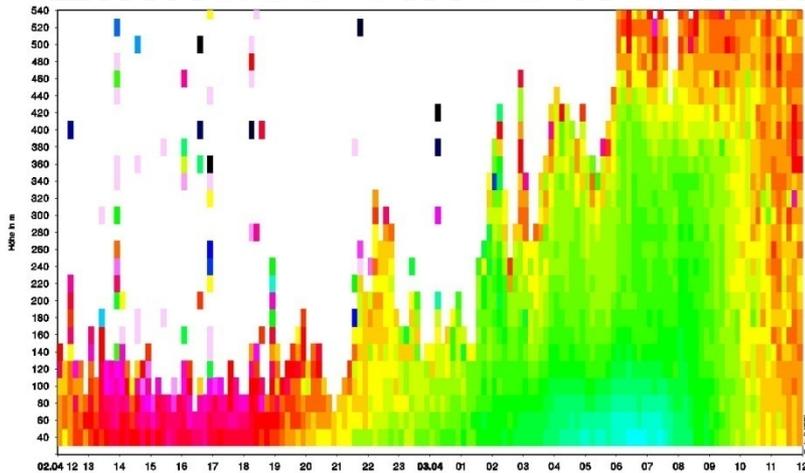


RASS

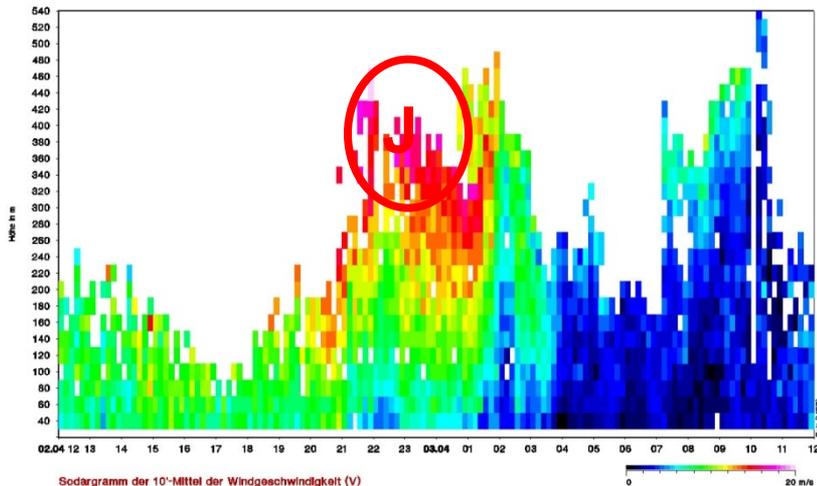


SODAR

Einfluss eines nächtlichen low-level jets



RASS pot. Temp. (max. ca. 25°C)

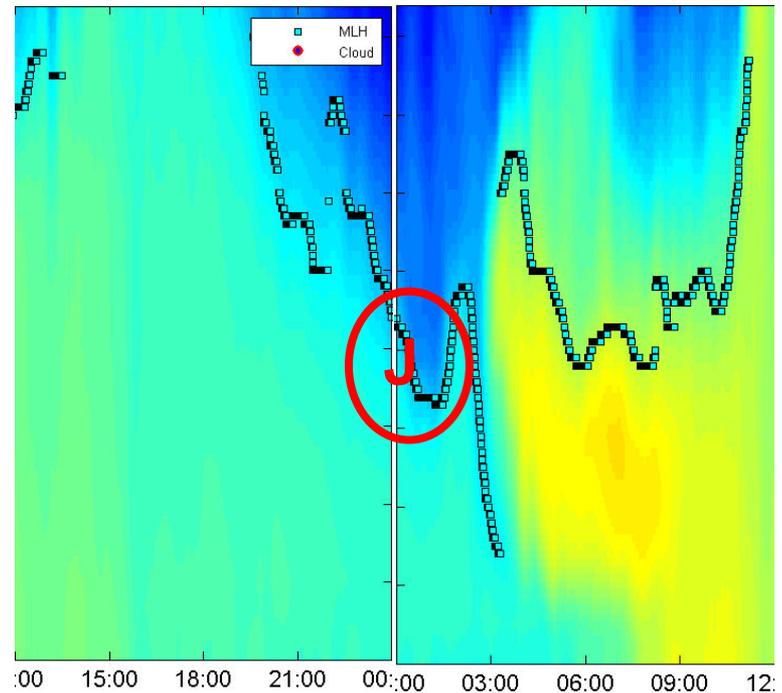


Sodarogramm der 10°-Mittel der Windgeschwindigkeit (V)

Augsburg

RASS (Sodar), mittlere horizontale Windgeschwindigkeit (max. 20 m/s)

825 m



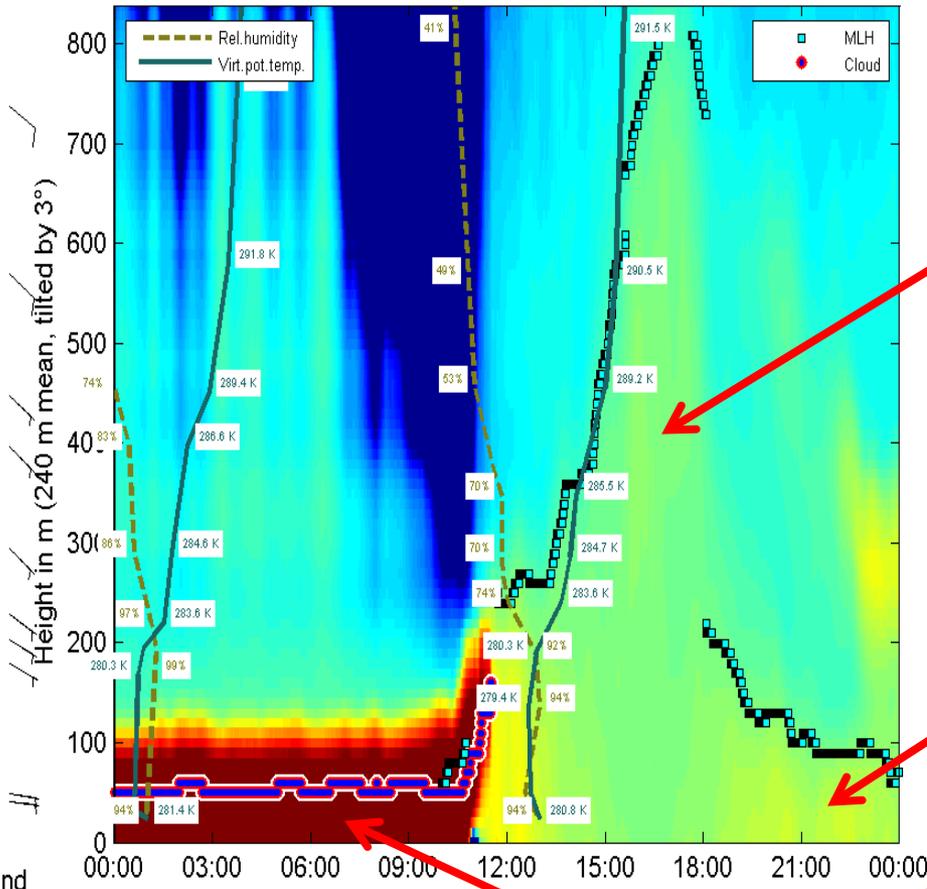
0 m

Ceilometer optische Rückstreuintensität (log. Skala)

Temperaturprofil und Luftqualität

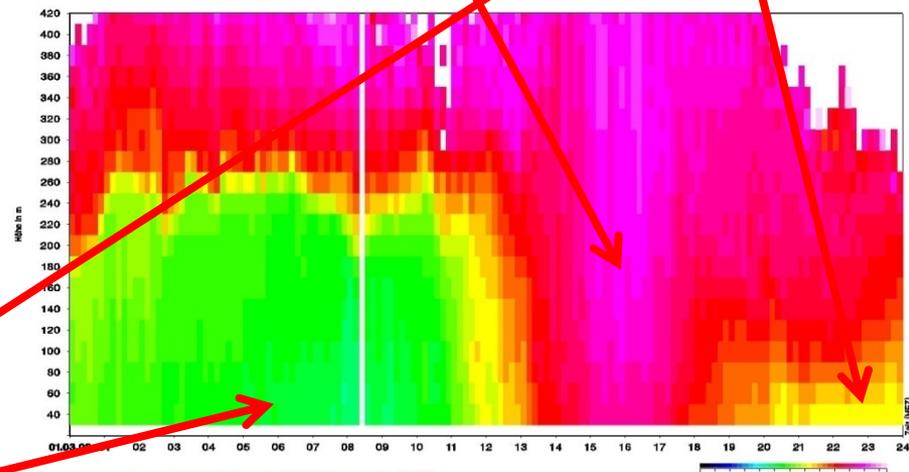
Aerosolrückstreuung (links), potenzielle Temperatur (rechts)

CL31 Augsburg AVA log₁₀ of backscatter with MLH on 01.03.2009 in 10⁻⁹ m⁻¹ sr⁻¹



gut durchmischt

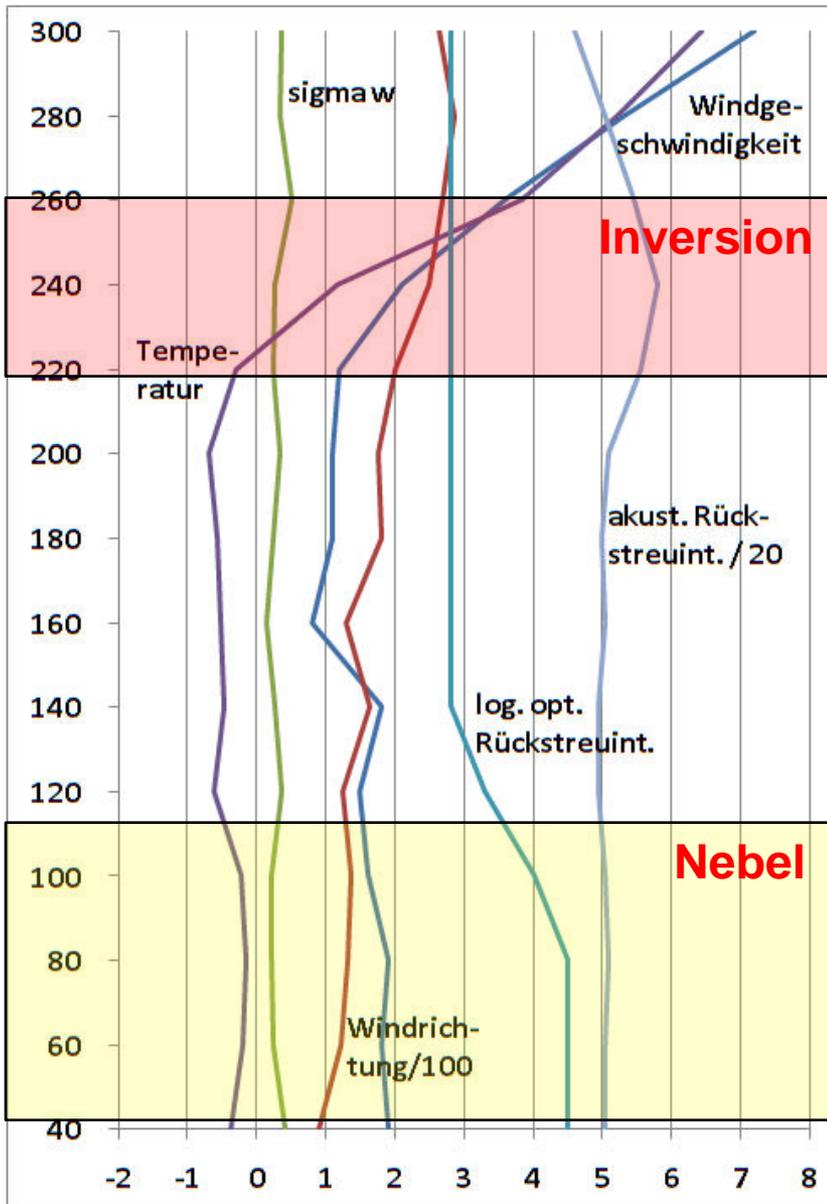
stabile nächtliche Grenzschicht



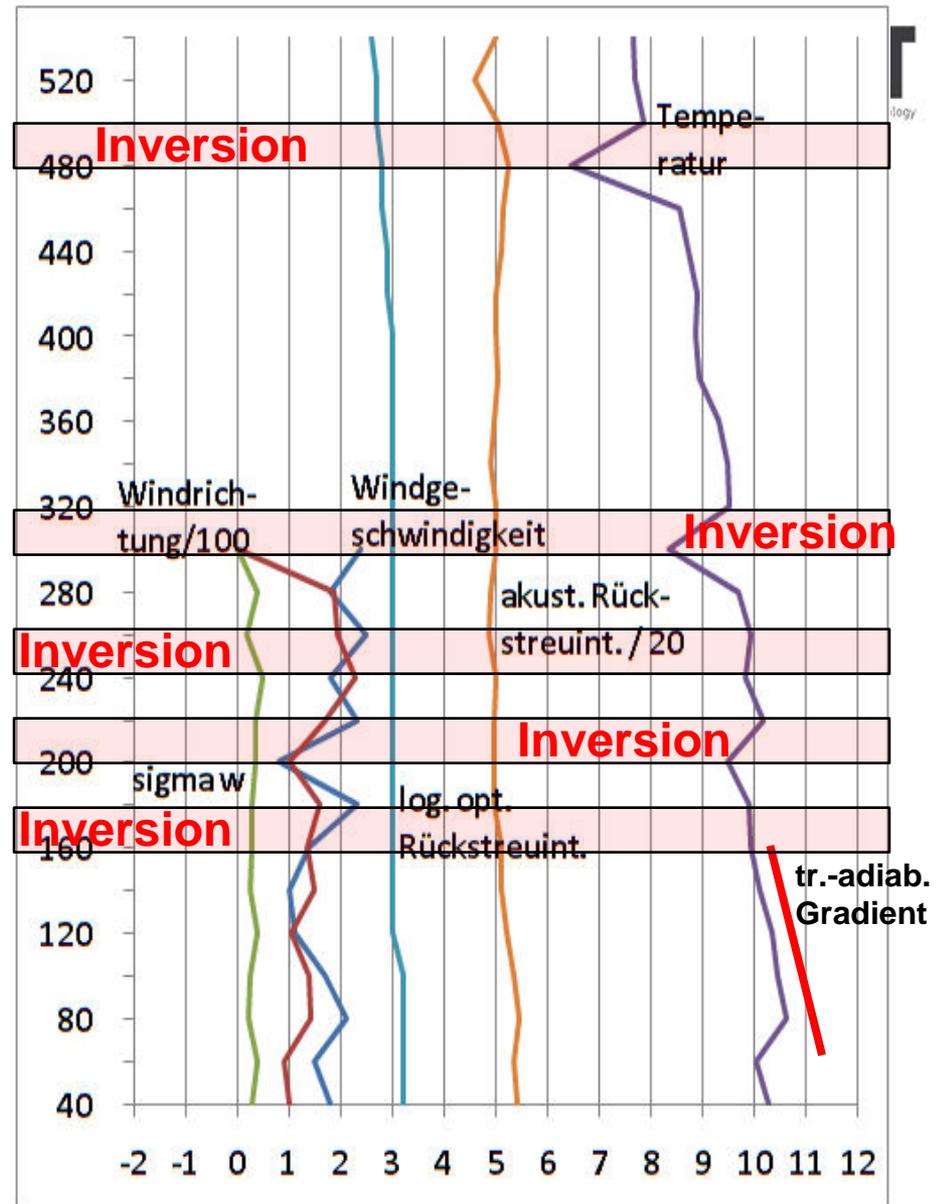
Sodargramm der 10'-Mittel der pot. RASS-Temperatur (PRT) Augsburg

RASS

Ceilometer **Nebel** 3.3XUEI MB 200



2009 03 01 06 Augsburg

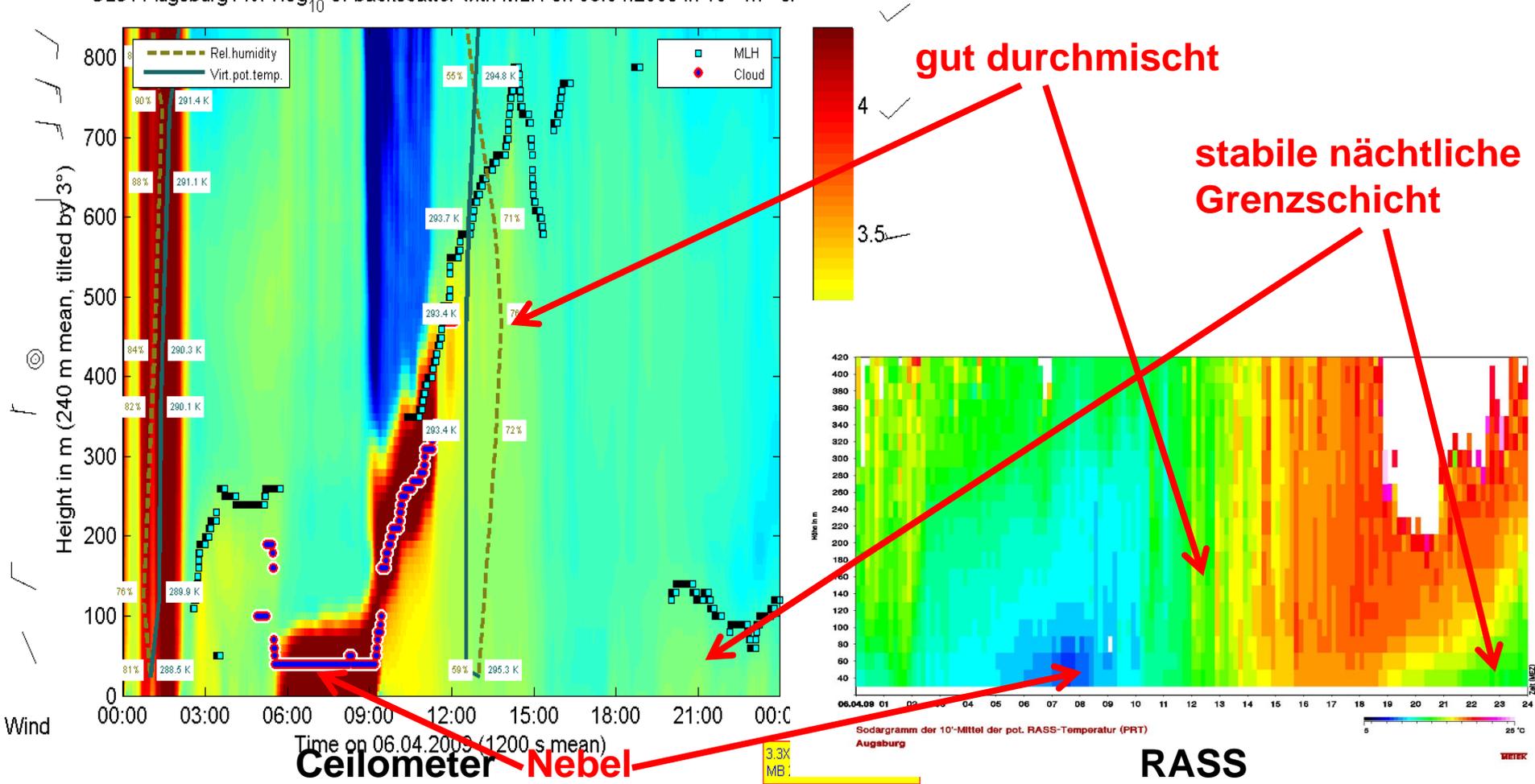


2009 03 01 15 Augsburg

Temperaturprofil und Luftqualität

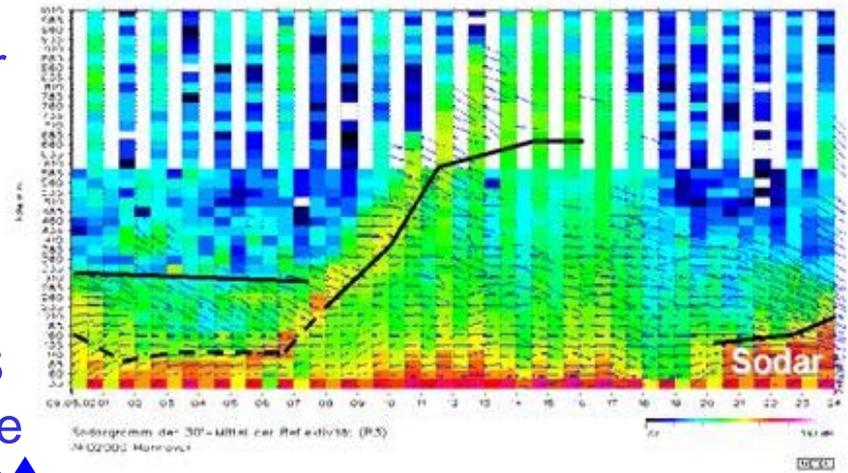
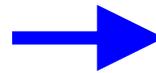
Aerosolrückstreuung (links), potenzielle Temperatur (rechts)

CL31 Augsburg AVA \log_{10} of backscatter with MLH on 06.04.2009 in $10^{-9} \text{ m}^{-1} \text{ sr}^{-1}$

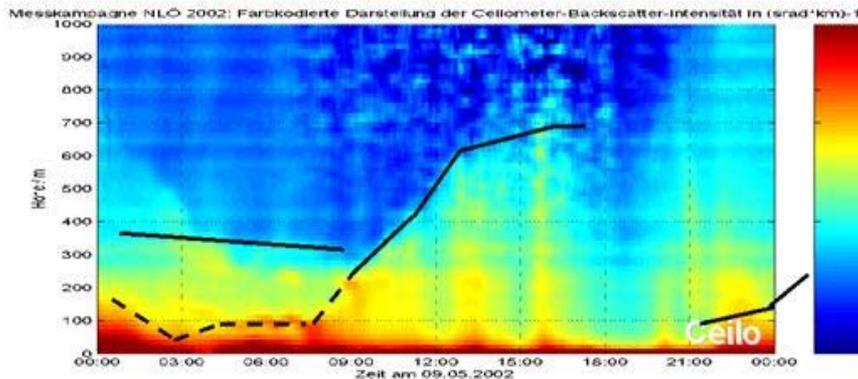


Comparison of MLH retrievals with three different remote sensing techniques

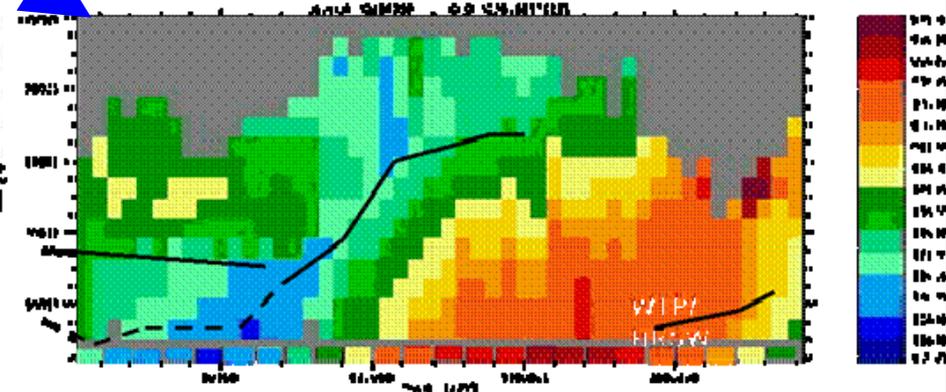
SODAR
acoustic backscatter



ceilometer
optical backscatter



RASS
temperature



Emeis, S., Chr. Münkel, S. Vogt, W.J. Müller, K. Schäfer, 2004: Atmospheric boundary-layer structure from simultaneous SODAR, RASS, and ceilometer measurements. Atmos. Environ., 38, 273-286.

Schlussfolgerungen

RASS-Sondierungen stützen die Interpretation von SODAR-Messungen

- Vertikale Windgradienten können über die Temperaturschichtung erklärt werden.
- Windsprünge können als Temperaturinversionen verifiziert werden.

RASS-Sondierungen stützen die Interpretation von Ceilometer-Messungen

- Obergrenzen hoher Rückstreuintensität können als Obergrenzen gut durchmischter Schichten identifiziert werden.
- Gebiete geringer Rückstreuintensität fallen gelegentlich mit Windmaxima (LLJ) zusammen.
- Einfluss der Luftfeuchtigkeit kann nicht mit RASS erfasst werden. Vergleiche mit Radiosondendaten deuten auch hier auf einen Zusammenhang hin.