

Description and parameterization of turbulence in the marine atmospheric boundary layer VERITAS TUFFO

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Gefördert auf Grund eines Beschlusses
des Deutschen Bundestages

Projekträger

Koordination

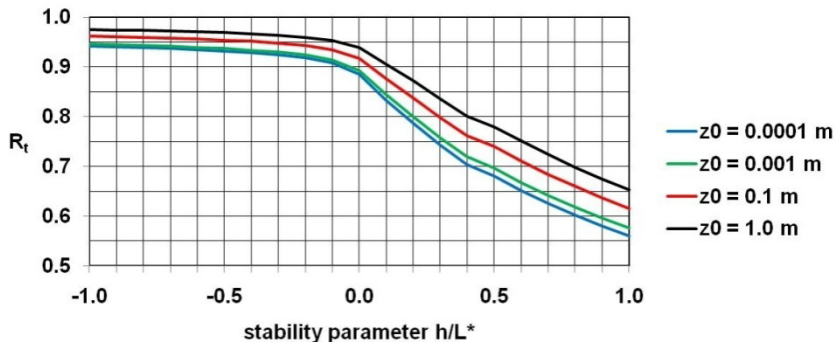
Meteorological challenges for offshore wind parks:

- marine atmospheric boundary layer is different:**
 - shallow atmospheric boundary layer**
 - wind speed-dependent roughness and turbulence**
 - wind direction and season-dependent atmospheric stability and turbulence**
 - stability-independent vertical gradient of atmospheric humidity**
- necessity for reliable wind field models with correct description/parameterization of turbulence**

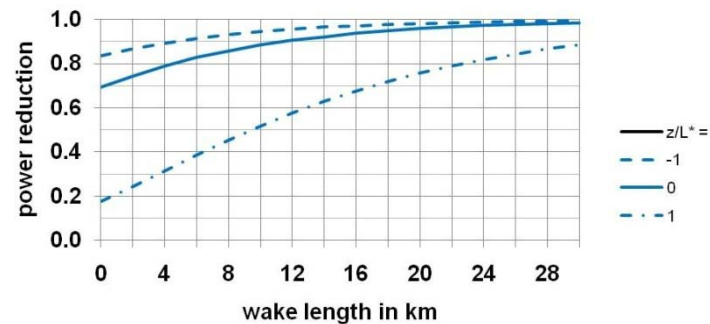
Turbulence

- influences loads on wind turbines (negative)
- influences harvests from wind turbines (positive)
- influences wind park efficiency (positive)
- influences wake lengths behind turbines and wind parks (positive)

power yield from large wind parks

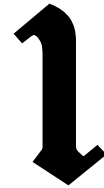


available power behind large wind parks



← more turbulence less turbulence →

Emeis, S., 2010: A simple analytical wind park model considering atmospheric stability. Wind Energy, 13, 459-469.



VERITAS results

VERITAS (July 1, 2008 to December 31, 2011)

Verification of **t**urbulence parameterization and description of the vertical structure of the maritime **a**tmospheric boundary layer in numerical **s**imulation models for wind analysis and forecast

became work package 5 of OWEA (see sessions 3 and 5 this morning)

investigators:

Richard Foreman M. Sc. (his PhD work, successfully completed Nov 31, 2011)
Prof. Dr. Stefan Emeis

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drag coefficient (neutral thermal stratification)

usual approach:

$$C_{Dn} = u_*^2 / U^2 \quad (\text{friction velocity over 10 m wind speed})$$

$$= \kappa^2 / \ln^2(z/z_0) \quad (\text{over land: logarithmic profile})$$



C_{Dn} is function of surface properties only

$$= \kappa^2 / \ln^2(gz/\alpha u_*^2) \quad (\text{over sea: Charnock's relation})$$



C_{Dn} depends on wind speed as well

empirically:

$$C_{D10n} = 0.000063 U_{10} + 0.00061 \quad (\text{Smith 1980})$$

no wave data available:

within the fully turbulent regime:

friction velocity:

$$u_* = 0.051 U_{10} - 0.14$$

input to the definition of C_D :

$$C_{D10n} = (0.051 U_{10} - 0.14)^2 / U_{10}^2$$

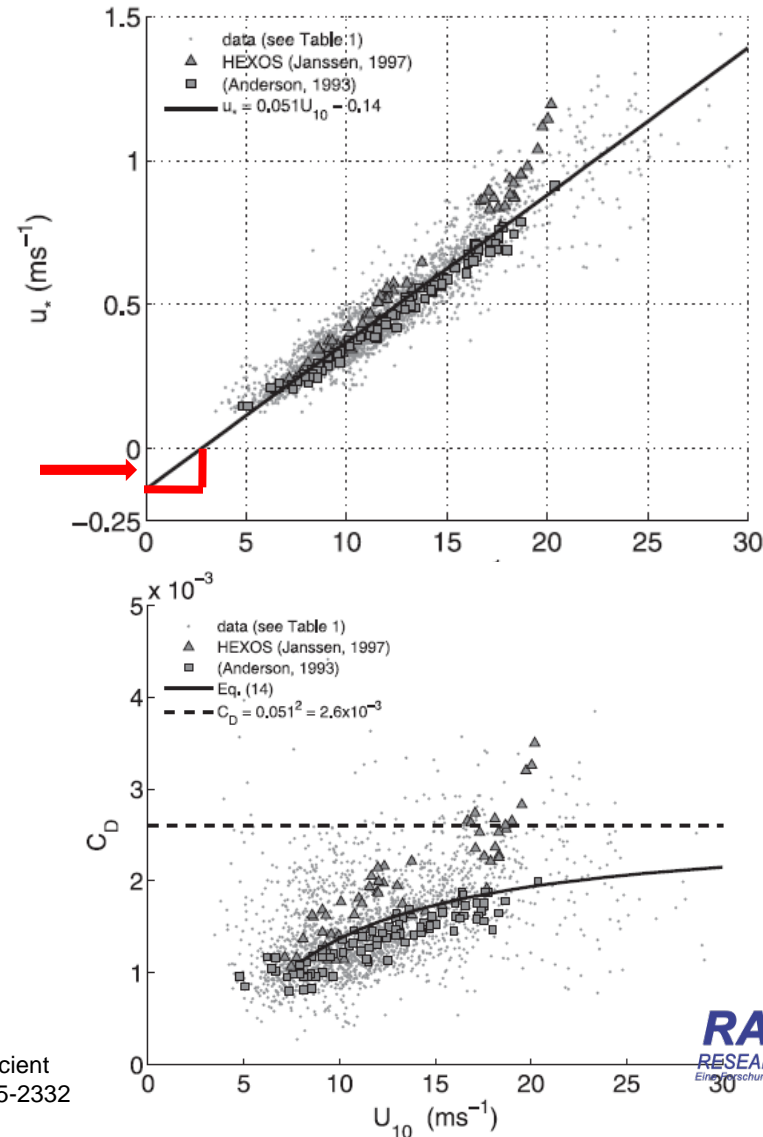
$$C_{D10n} = (C_m U_{10} - b)^2 / U_{10}^2$$

approaches:

$$C_m^2 = 0.051^2 = 0.0026$$

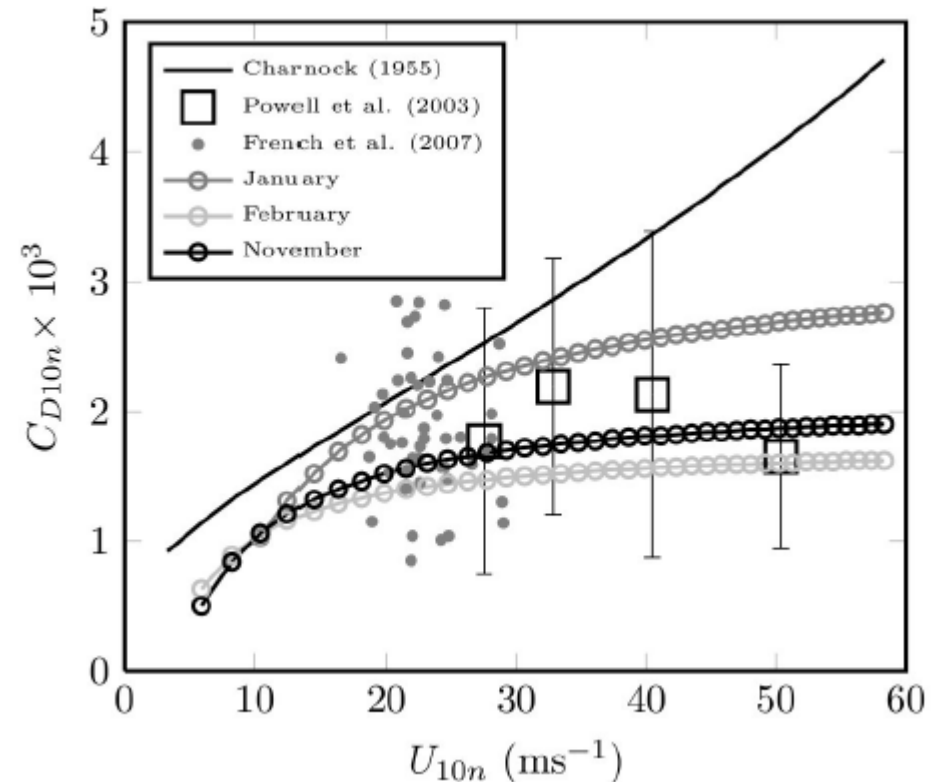
for large U_{10}

Foreman, R., S. Emeis, 2010: Revisiting the Definition of the Drag Coefficient in the Marine Atmospheric Boundary Layer. J. Phys. Oceanogr., **40**, 2325-2332



Comparison with data:

	C_m	b
Literatur	0.051	0.14 m/s
FINO Jan 2005	0.057	0.26 m/s
FINO Feb 2005	0.042	0.01 m/s
FINO Nov 2005	0.048	0.02 m/s



🌪️ C_{D10n} still depends on wind speed (but at least: C_m does not)
the plot shows a non-dimensional variable plotted against a dimensional one

wave data available:

$$C_{D10n} = a (H_s/\lambda_p)^2 \quad \text{(from dimensional analysis)}$$

the drag coefficient should not approach zero for vanishing waves. Therefore a minimum is set:

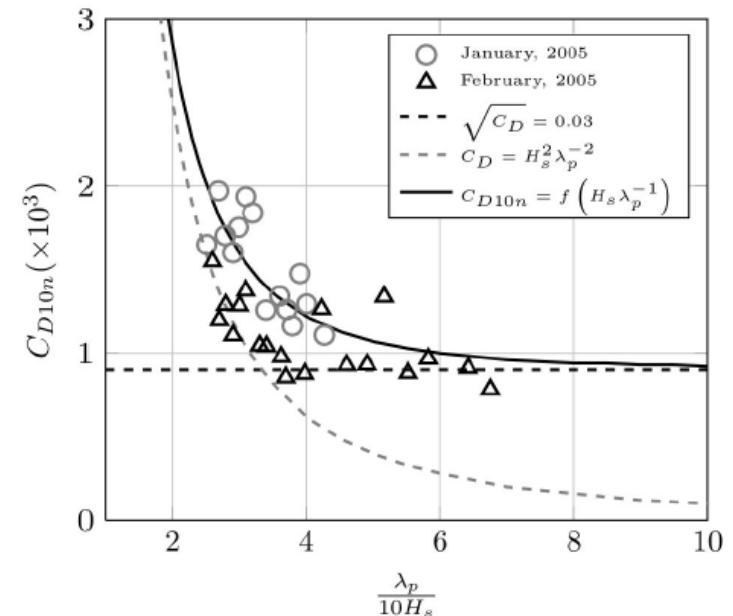
$$C_{D10n} = 0.0009 \quad \text{(smooth surface)}$$

fit between both extremes:

$$(C_{D10n}^{1/2})^n = 0.03^n + (H_s/\lambda_p)^n$$

empirically $n \approx 3$

(Churchill and Usagi 1972)



steep ← → shallow

- ☺ C_{D10n} depends on surface parameter only
- ☺ both axes in plot are now non-dimensional

**turbulence parameterization in meso-scale wind field models
such as MM5 or WRF:**

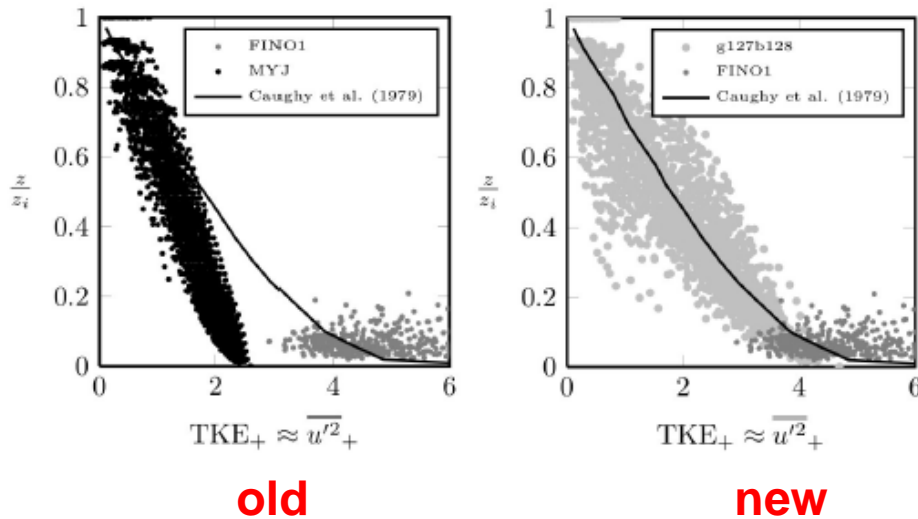
	A1	A2	B1	B2	C1
MY 1982	0.92	0.74	16.6	10.1	0.08
MYJ 2002	0.660	0.657	11.878	7.227	0.00083
new	0.91	0.54	28.76	13.08	0.15

plus an adaptation of turbulent length scales following Nakanishi (2001)

new values are based on modern laboratory data
- at very high Reynolds numbers
- with very small velocity sensors close to the wall

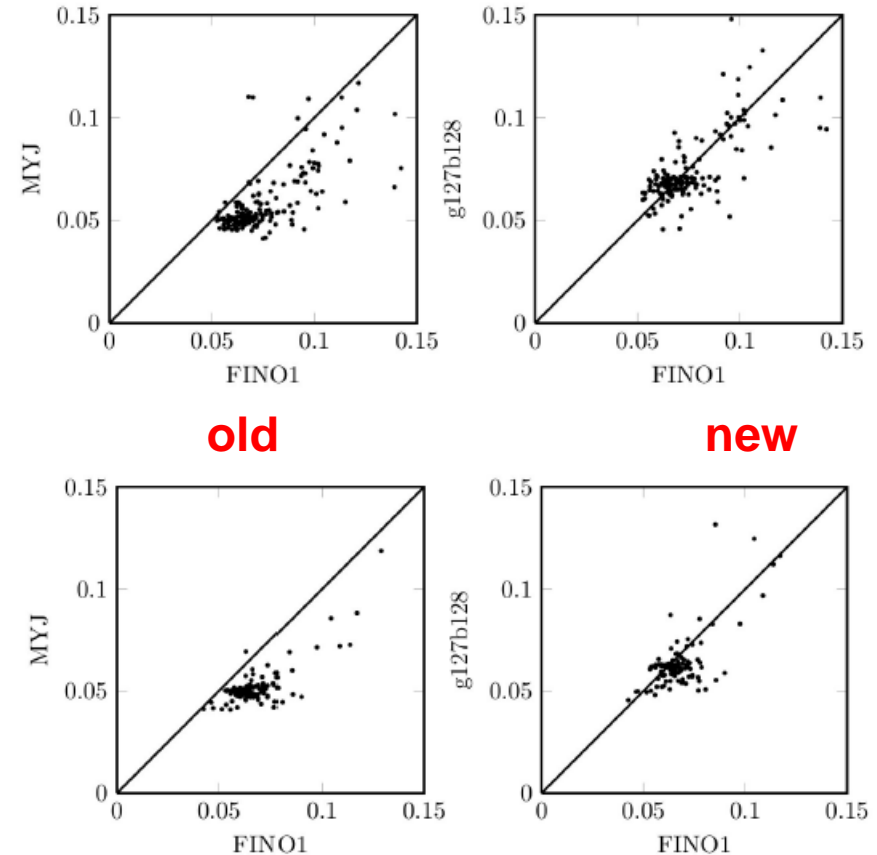
Foreman, R., S. Emeis, 2012: Method for increasing Turbulent Kinetic Energy in the Mellor-Yamada-Janjić boundary layer parametrization. accepted by Bound.-Lay. Meteorol.

comparison of model results with offshore (FINO1) and onshore (Caughey et al. 1979) data



vertical profiles of
normalised turb. kin. energy
January 2005

turbulence intensity at 80 m height February 2005



turbulence intensity at
80 m height November 2005

2

TUFFO idea

TUFFO (August 1, 2011 to July 31, 2014)

Detection and assessment of the impact of **turbulent** humidity (**F**euchte) **f**luxes on turbulence in **o**ffshore wind parks

investigators:

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humid air is lighter than dry air

sea surface is a perfect humidity source

- ➔ near surface air is nearly saturated
- ➔ air aloft is less saturated

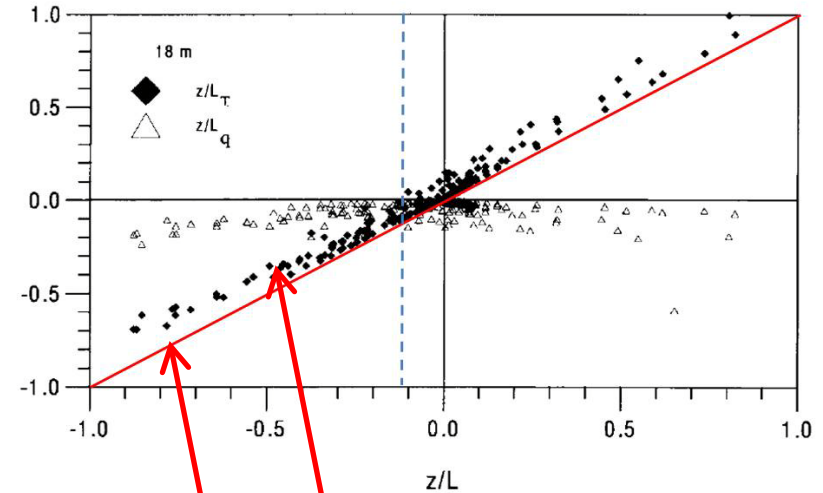
➔ less static stability (in 30 to 50 % of all cases: humidity profile decides on static stability)

(Sempreviva and Gryning 1996, Edson et al. 2004)

➔ more turbulence

➔ more loads on turbines

➔ less wake lengths behind turbines and wind parks



stability from
temperature only
true stability of air mass

new approach in TUFFO:

**deployment of high-speed humidity sensors to FINO 1
co-located with ultra-sonic anemometers**

- high-resolution humidity data**
- turbulent vertical humidity fluxes**
- better static stability information**

assessment of impact on turbulence at hub height

assessment of impact on wind parks (efficiency, wake lengths)

**update of turbulence parameterization in meso-scale wind field
models**

Summary:

VERITAS

sea surface drag description

no wave data: sea surface drag flattens off for high wind speeds

with wave data: sea surface drag depends on wave steepness squared

meso-scale wind field models

enhanced turbulence parameterization which gives higher (more realistic) turbulence intensities in the lower part of the atmospheric boundary layer

TUFFO

vertical humidity structure in the marine boundary layer

leads to more unstable static stratification → more turbulence

Thank you very much for your attention

Gefördert auf Grund eines Beschlusses
des Deutschen Bundestages



Bundesministerium
für Umwelt, Naturschutz
und Reaktorsicherheit

Projekträger

