

Development and application of a coupled atmospheric-hydrological model system, suitable for regional spatial and climate relevant temporal scales


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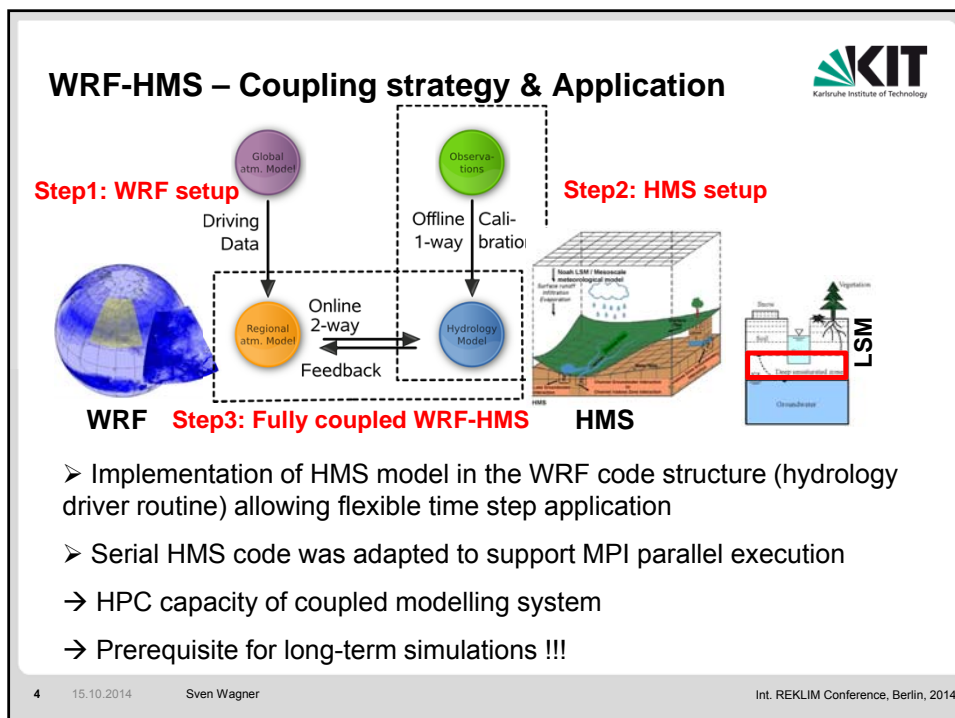
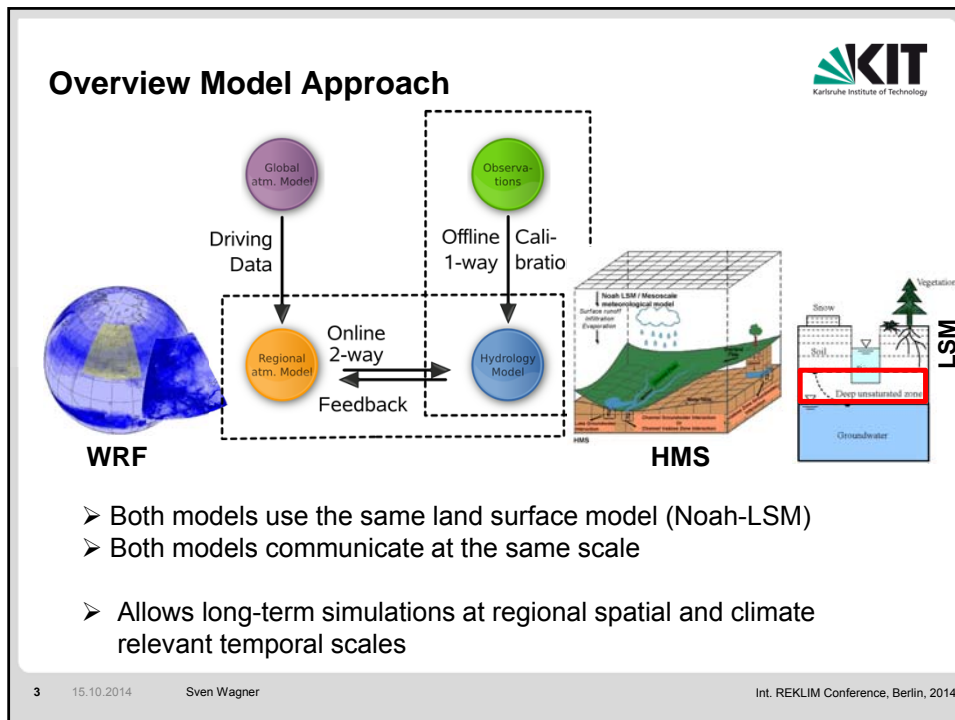
Objectives

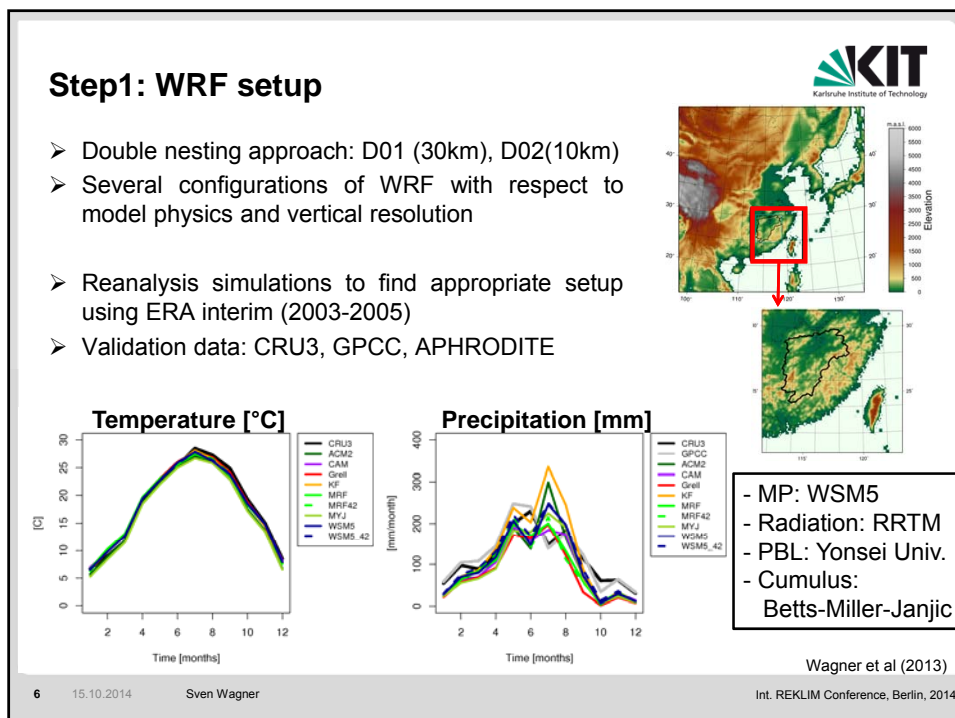
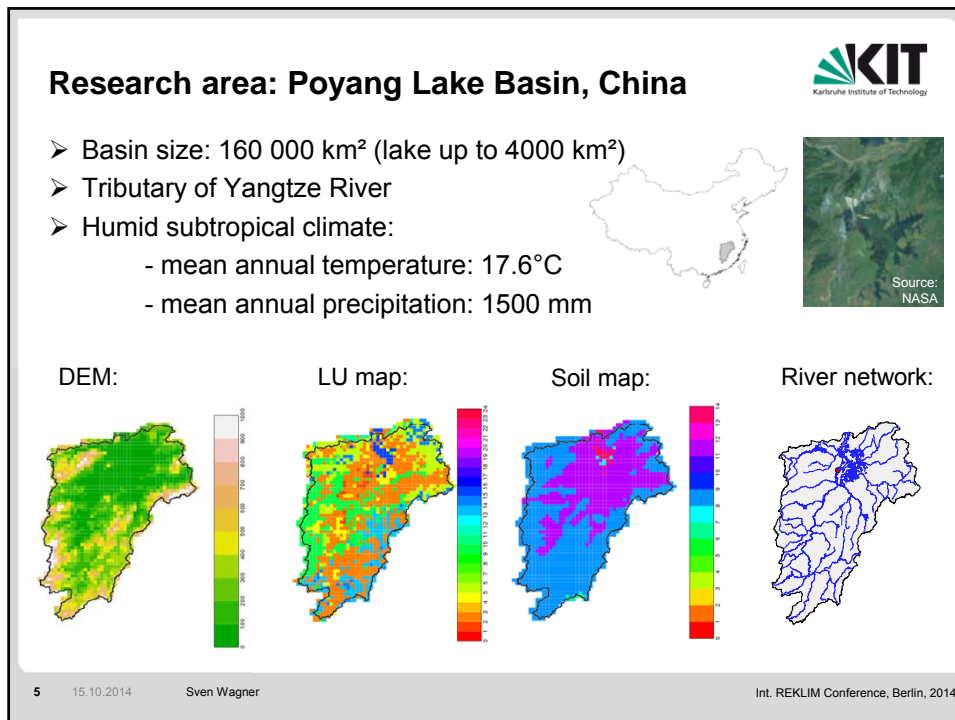
- All components of hydrological cycle are affected by climate and landuse changes
- Joint landuse- & climate change impact analysis on regional water cycle requires

Investigations on Feedback mechanisms between the atmosphere, land surface & subsurface conditions


- The quantification of such feedback mechanisms calls for coupled modelling systems that consist of a
 - regional atmospheric- &
 - distributed hydrological model
 - sharing compatible water & energy flux formulations

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Step2: NOAH-LSM – HMS simulations




- Meteorological forcing: interpolated station data
- use implemented HMS model in the WRF code
 - same modules & input data (except met. forcing)

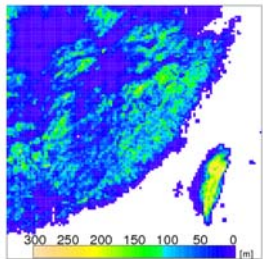


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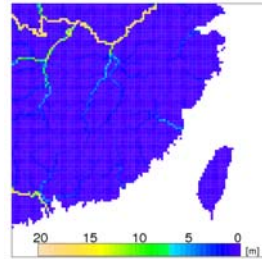
Step 2: NoahLSM-HMS – PREPROCESSING: Additional hydrological input parameters



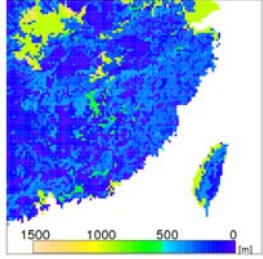
DEM (sd):
USGS
HYDRO1K
(GTOPO30)



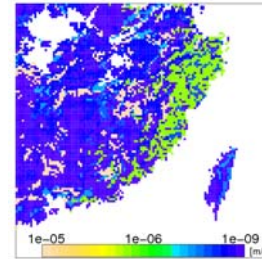
Streambed depth:
USGS
HYDRO1K



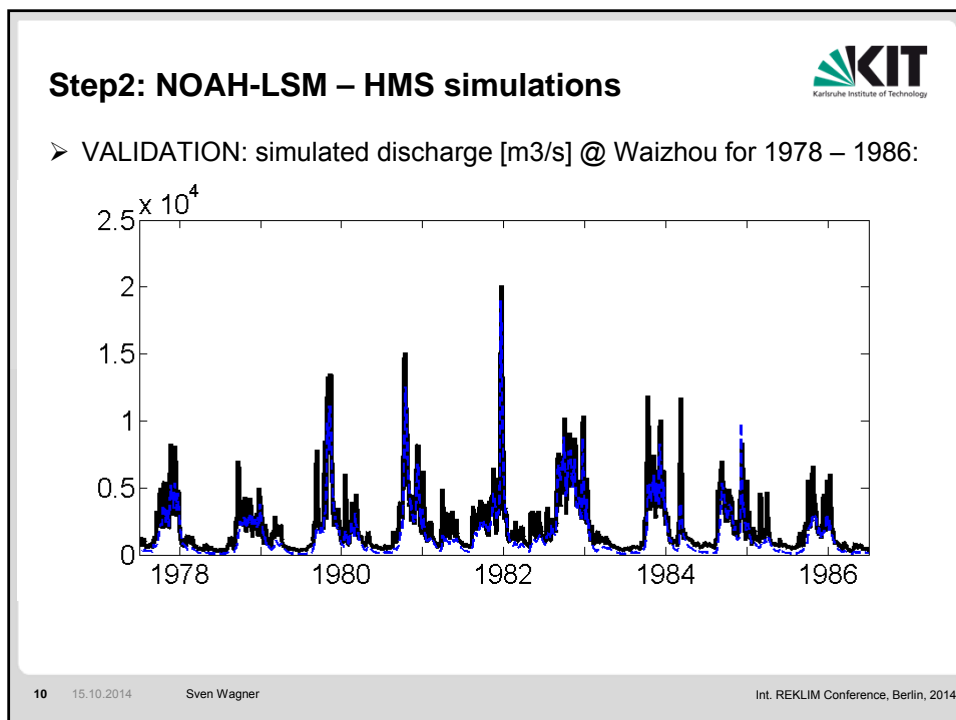
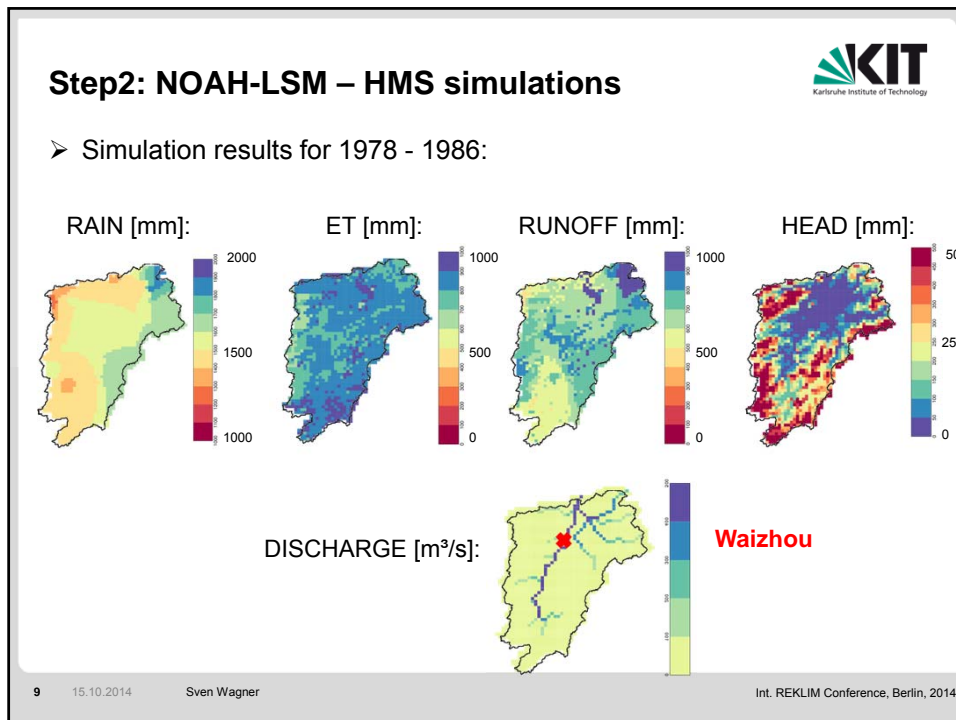
Aquifer thickness:
Chinese
Geological
data set




Hydraulic conductivity:
Chinese
Geological
data set



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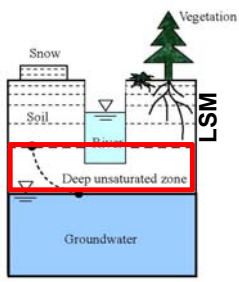
Step2: NOAH-LSM – HMS with GW-feedback



Methods for allowing feedbacks between LSM and saturated zone


- Coupling of saturated to unsaturated zone
- Two way interaction & fluxes (e.g. capillary rise vs. gravity fluxes) between saturated and unsaturated zone

- **Richard's equation with *fixed-head boundary condition*** based on Zeng et al. (2009), De Rooij (2010)
- **Darcy flux boundary condition** based on Bogaart et al. (2008)



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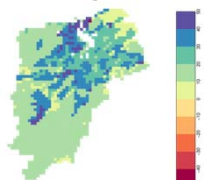
Step2: NOAH-LSM – HMS with GW-feedback




Impact of feedback between LSM and saturated zone on simulation results

- Difference plots “Fixed head” versus “no Coupling”

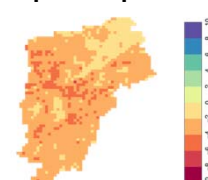
Recharge [%]



Soil moisture 0–10 cm

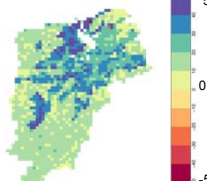


Evapotranspiration [%]

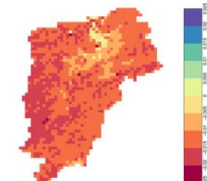


- Difference plots “Darcy-Flux” versus “no Coupling”

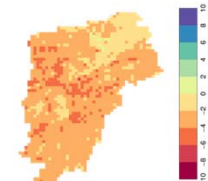
Recharge [%]



Soil moisture 0–10 cm




Evapotranspiration [%]

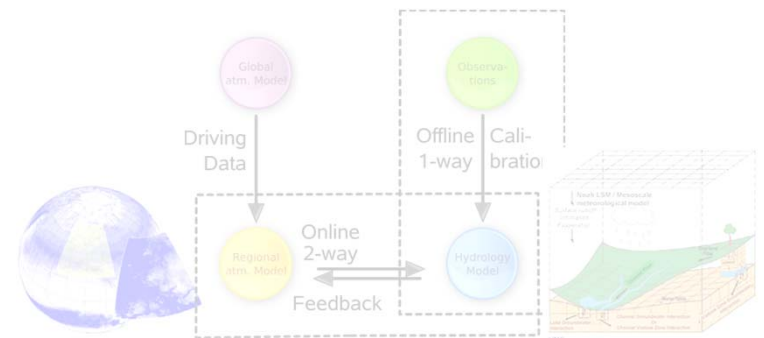


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Step3: WRF – NOAH-LSM – HMS simulations




- Use the identified optimal stand-alone WRF and HMS setup
- Allows investigations of hydrological land surface – atmosphere feedback



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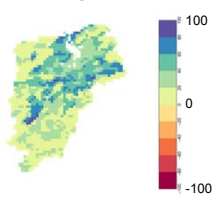
Step3: WRF – NOAH-LSM – HMS simulations



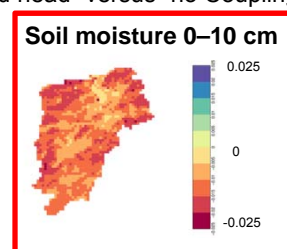
Fully coupled simulation results with GW feedback

- Difference plots “Fixed head” versus “no Coupling”

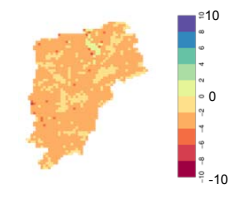
Recharge [%]



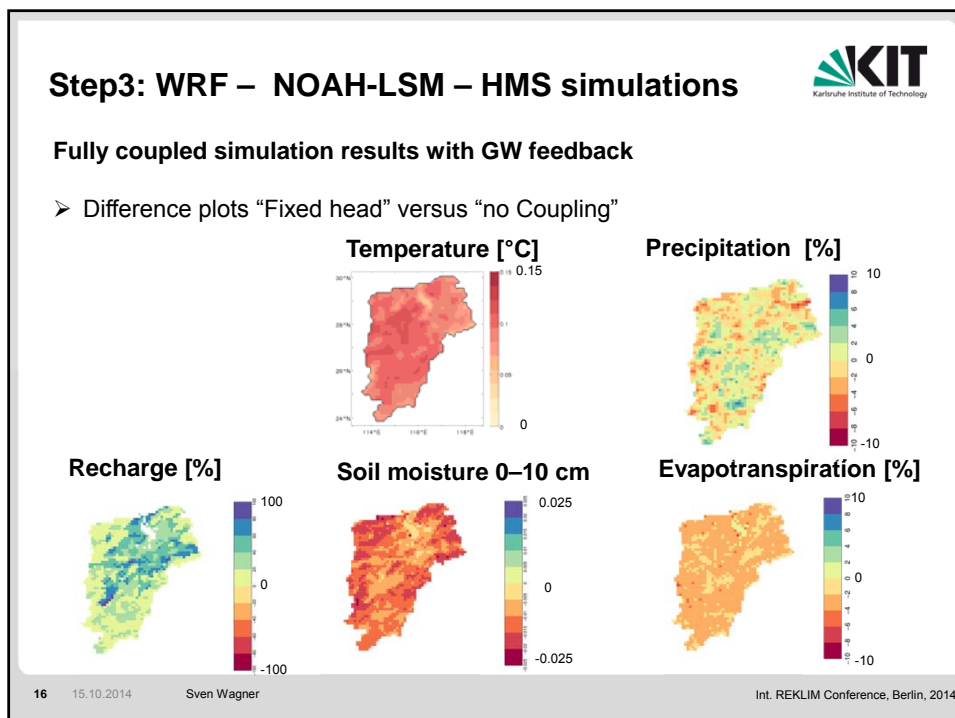
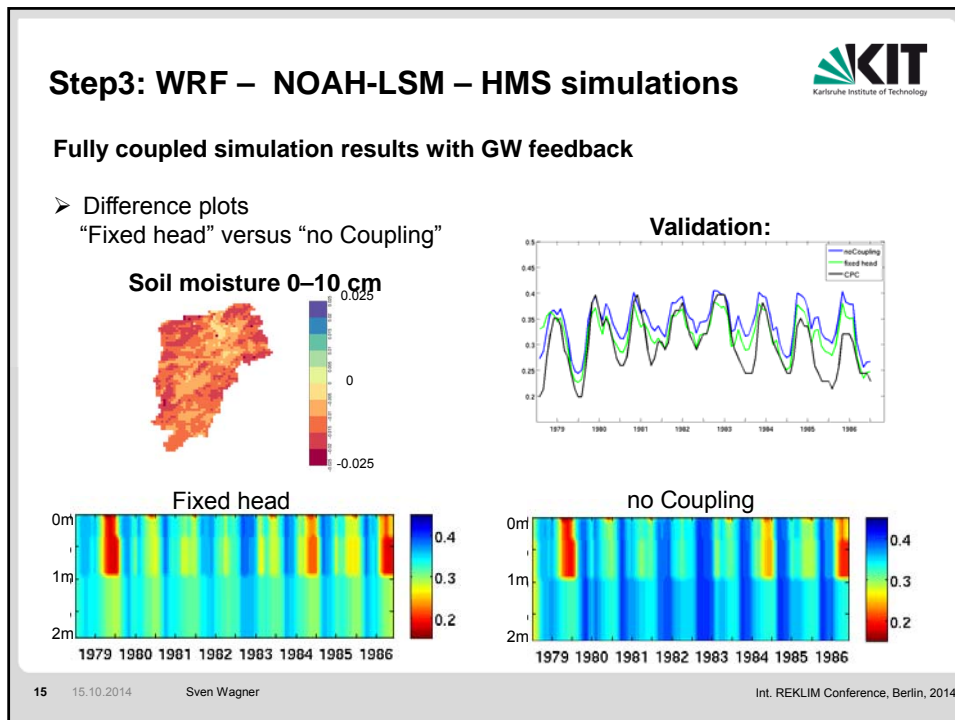
Soil moisture 0–10 cm



Evapotranspiration [%]



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Summary



Objective: Investigations on Feedback mechanisms between the atmosphere, land surface & subsurface conditions

Fully coupled modelling system:

- Integration of HMS preprocessors & code in WRF model structure
- Integration of GW feedback mechanisms in coupled model system

Poyang Lake Basin:

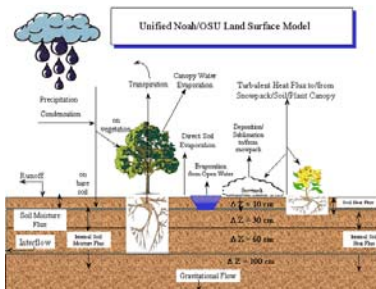
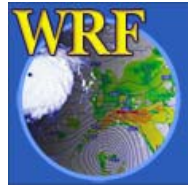
- Performance & potential of fully coupled simulations incl. GW-feedback
- GW-feedback impact conditions in LSM, land surface & lower atmosphere

Thank you for your attention

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Model Components: WRF & Noah-LSM



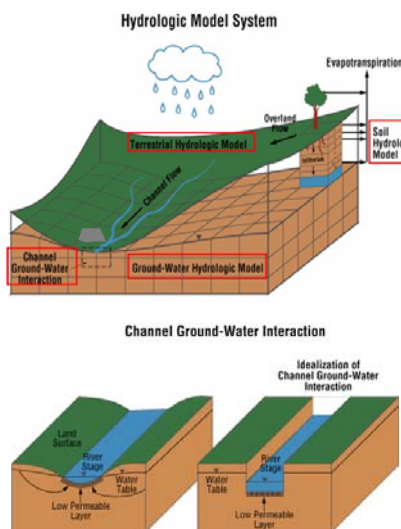
- Non-hydrostatic, $\Delta x \approx 1$ km-50 km, $\Delta t \approx$ tens of seconds
 - Based on conservation laws
 - Subgridscale processes: parameterized
 - Nested approach: lateral boundary from GCM
 - “Lower boundary of WRF”
 - $\Delta t \approx$ tens of minutes
 - 4 soil layers
 - Vertical water and energy fluxes
- Important for feedbacks between near surface soil, boundary layer & atmosphere

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Model Components: HMS (Yu et al, 2006)



- Spatially distributed
- Suitable for large- scale applications , $\Delta x \approx$ up to a few tens of km: 10 km here
- streamflow routing - 2D diffusive wave, $\Delta t \approx$ tens of minutes
- Interaction of channel & vadose zone or channel & groundwater flux
- Unsaturated soil moisture profile is assumed to be in equilibrium
- 2D horizontal groundwater flow: one layer aquifer, simple bedrock, $\Delta t \approx 1$ day


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Step2: NOAH-LSM – HMS simulations

- Meteorological forcing: interpolated station data
- use implemented HMS model in the WRF code
 - same modules & input data (except met. forcing)
- Calibration of HMS:
 - Noah LSM parameters: FXEXP, REFKDT, ...
 - HMS parameters: Manning, conductivities, ...






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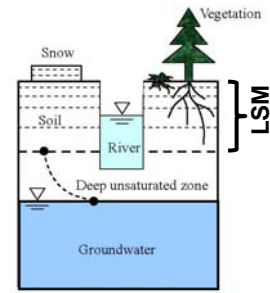
Feedback between LSM and saturated zone

Approach 1: Richard's equation with *fixed-head boundary condition*

Based on Zeng *et al.* (2009), De Rooij (2010)

- Free drainage boundary condition of LSM is replaced by a **fixed-head bottom boundary condition** which assumes an equilibrium soil moisture distribution
- Hydraulic head & soil moisture at the lower boundary of LSM is derived from **distance between groundwater level and bottom of LSM** conserving the energy and mass of water
- New boundary condition realized with **additional layer** at bottom of Noah-LSM
- Label: **Fixed-Head**





Distinction of 3 cases: new layer is un-, partially- or fully saturated, e.g.

$$\bar{\theta}_i = \frac{\int_{z_{i-1/2}}^{z_{i+1/2}} \theta_R(z) dz}{z_{i+1/2} - z_{i-1/2}} = \frac{\theta_{sat} \psi_{sat}}{(z_{i+1/2} - z_{i-1/2}) \left(\frac{1}{B} - 1 \right)} \times \left[\left(\frac{\psi_{sat}}{z_p - z_{i+1/2}} \right)^{1/B-1} - \left(\frac{\psi_{sat}}{z_p - z_{i-1/2}} \right)^{1/B-1} \right]$$

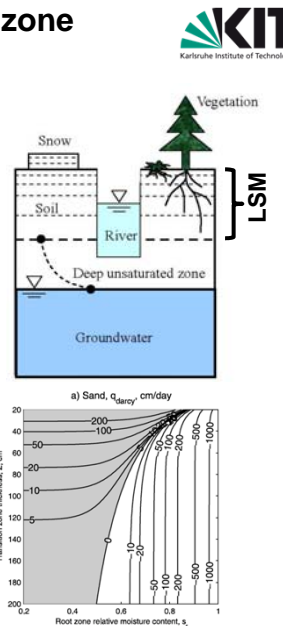
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Feedback between LSM and saturated zone

Approach 2: Darcy flux boundary condition

Based on *Bogaart et al. (2008)*

- Assumes a quasi steady-state moisture profile between groundwater head and lowest soil layer of the LSM.
- Darcy equation** is used to describe flow through this transition zone depending on relative saturation at bottom of LSM
- Parameterization** that approximates net Darcy flux q_{darcy} for different thicknesses of transition zone and different values of saturation for lowest LSM soil layer
- Label: **Darcy-Flux**



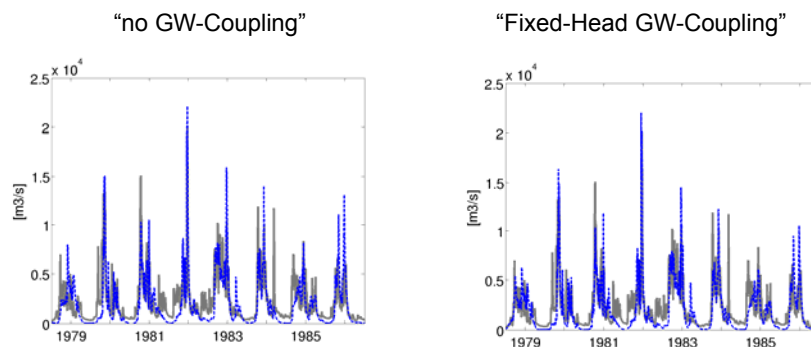
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Step3: WRF – NOAH-LSM – HMS simulations

➤ VALIDATION: simulated discharge [m³/s] @ Waizhou for 1978 – 1986 of fully coupled model system



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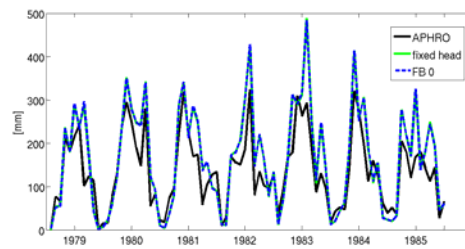
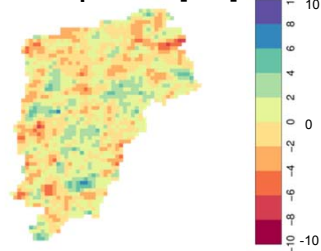
Step3: WRF – NOAH-LSM – HMS simulations



Fully coupled simulation results with GW feedback

- In fully coupled mode: in addition impact on atmospheric variables
- e.g. **Precipitation**
- Difference plots “Fixed head” versus “no Coupling”

Precipitation [mm]



WRF-NoahLSM-HMS – Coupling strategy:



- Implementation of HMS model in the WRF code structure (hydrology driver routine) allowing flexible time step application
 - Integration of preprocessors (static surface and sub-surface hydrological parameters)
 - NetCDF compliance (IO)
 - Serial HMS code was adapted to support MPI parallel execution
- HPC capacity of coupled modelling system
- Prerequisite for long-term simulations !!!

