

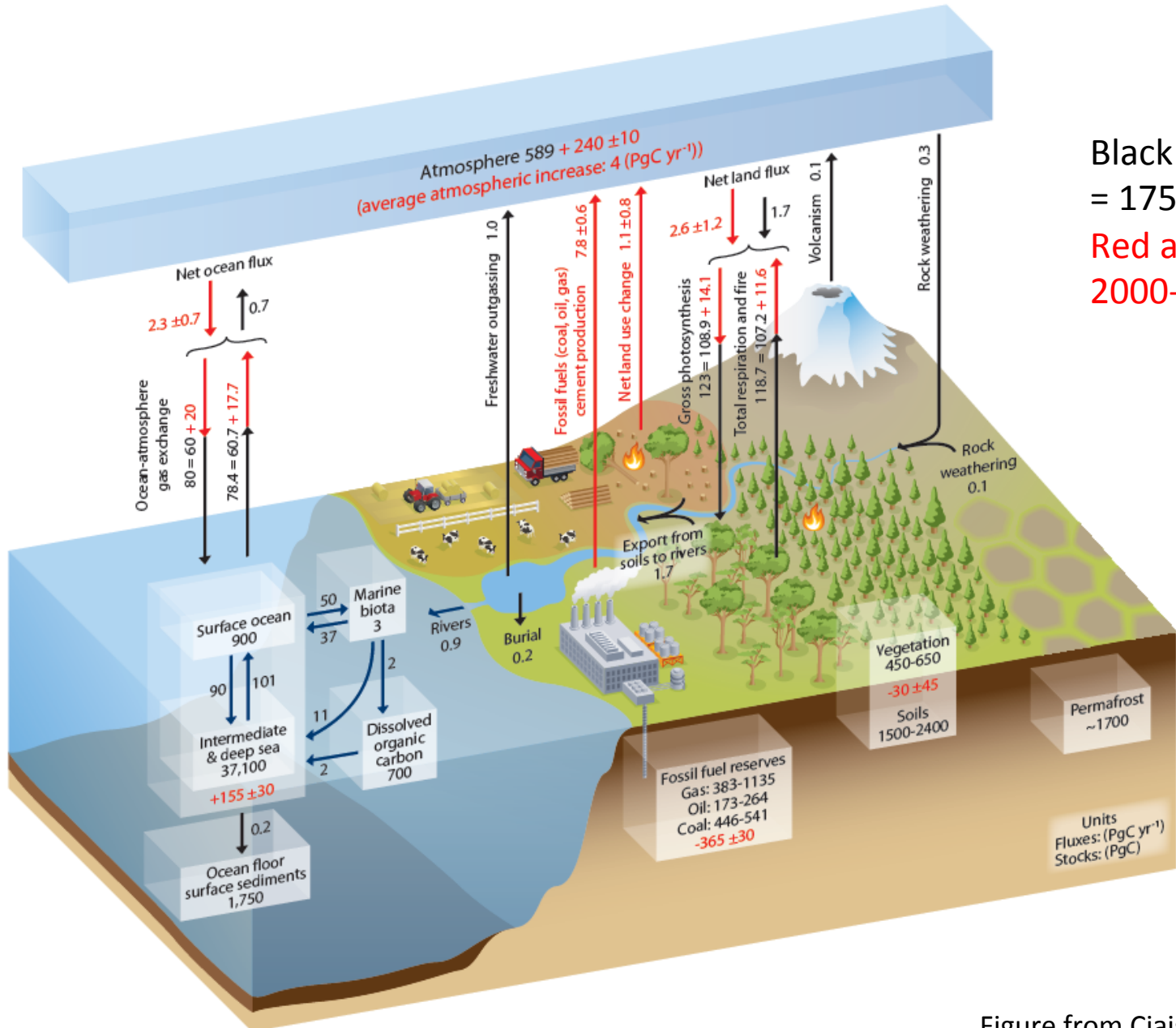
An overview of the terrestrial carbon-cycle from a global modelling perspective

Tom Pugh

Karlsruhe Institute of Technology, IMK-IFU, 82467 Garmisch-Partenkirchen, Germany.



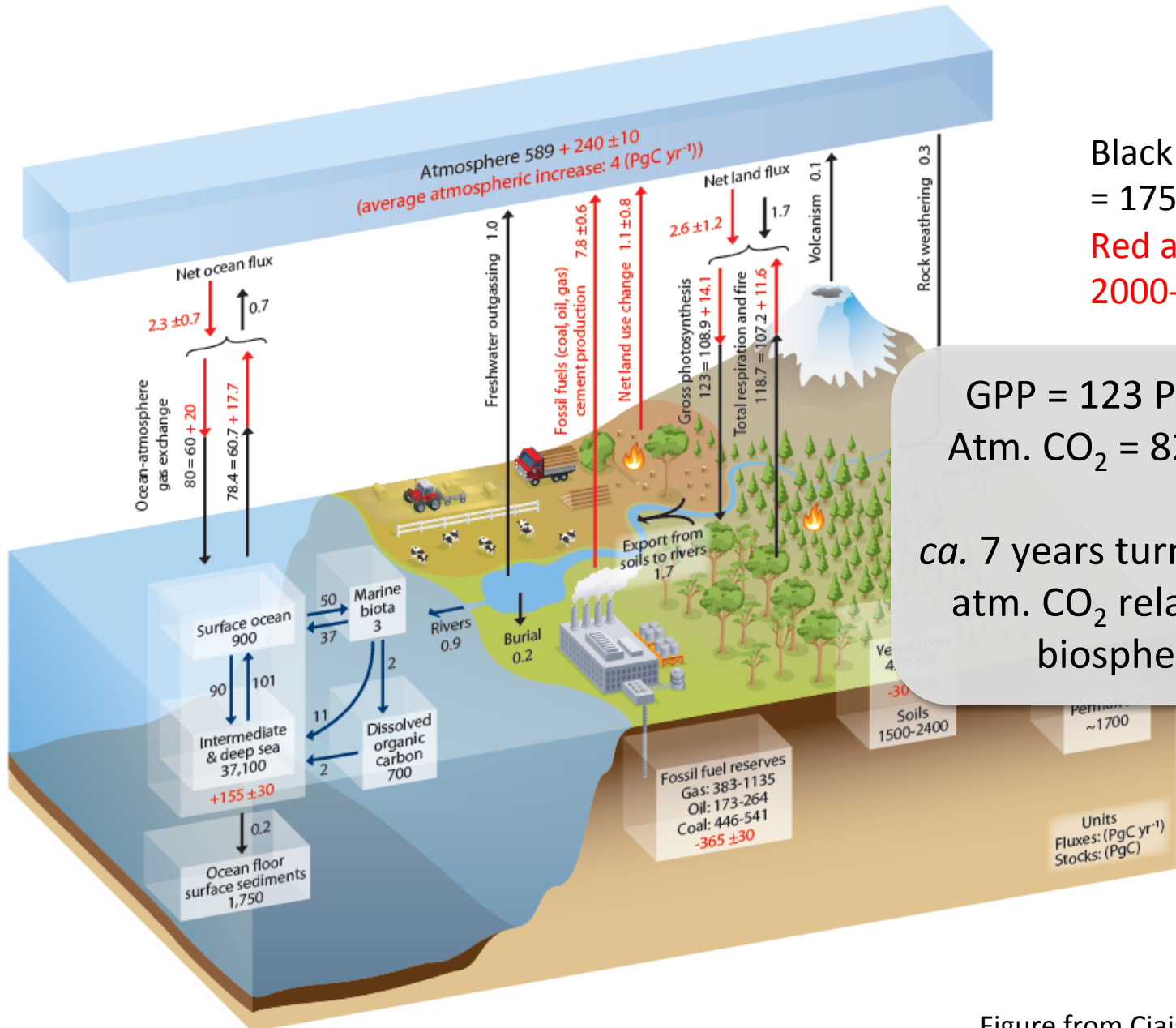
Global carbon-cycle overview



Black arrows = 1750
Red arrows = 2000-2009

Figure from Ciais et al. (2013)

Global carbon-cycle overview

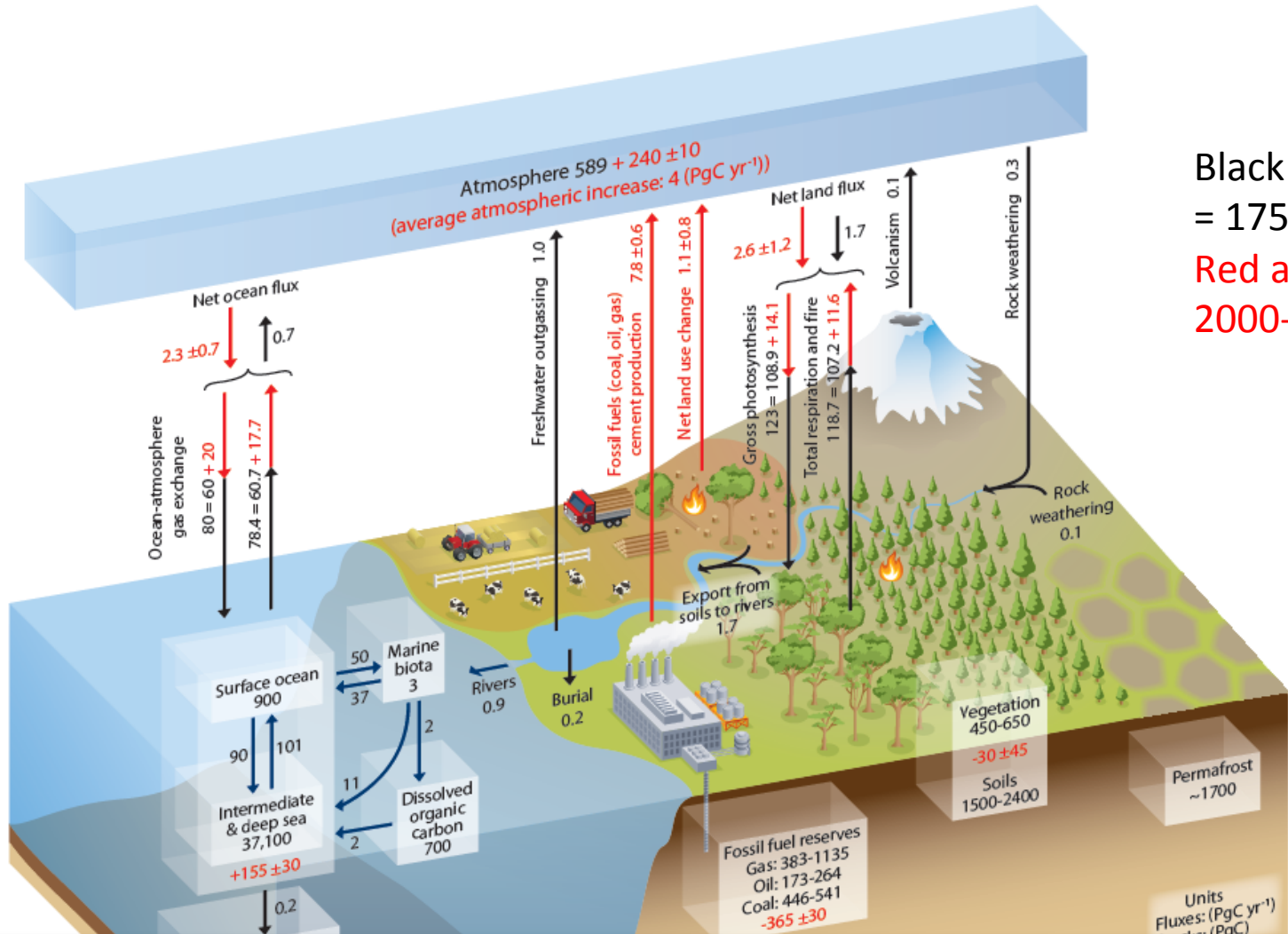


Black arrows = 1750
Red arrows = 2000-2009

GPP = 123 Pg C a⁻¹
 Atm. CO₂ = 829 Pg C
 ca. 7 years turnover for atm. CO₂ relative to biosphere

Figure from Ciais et al. (2013)

Global carbon-cycle overview

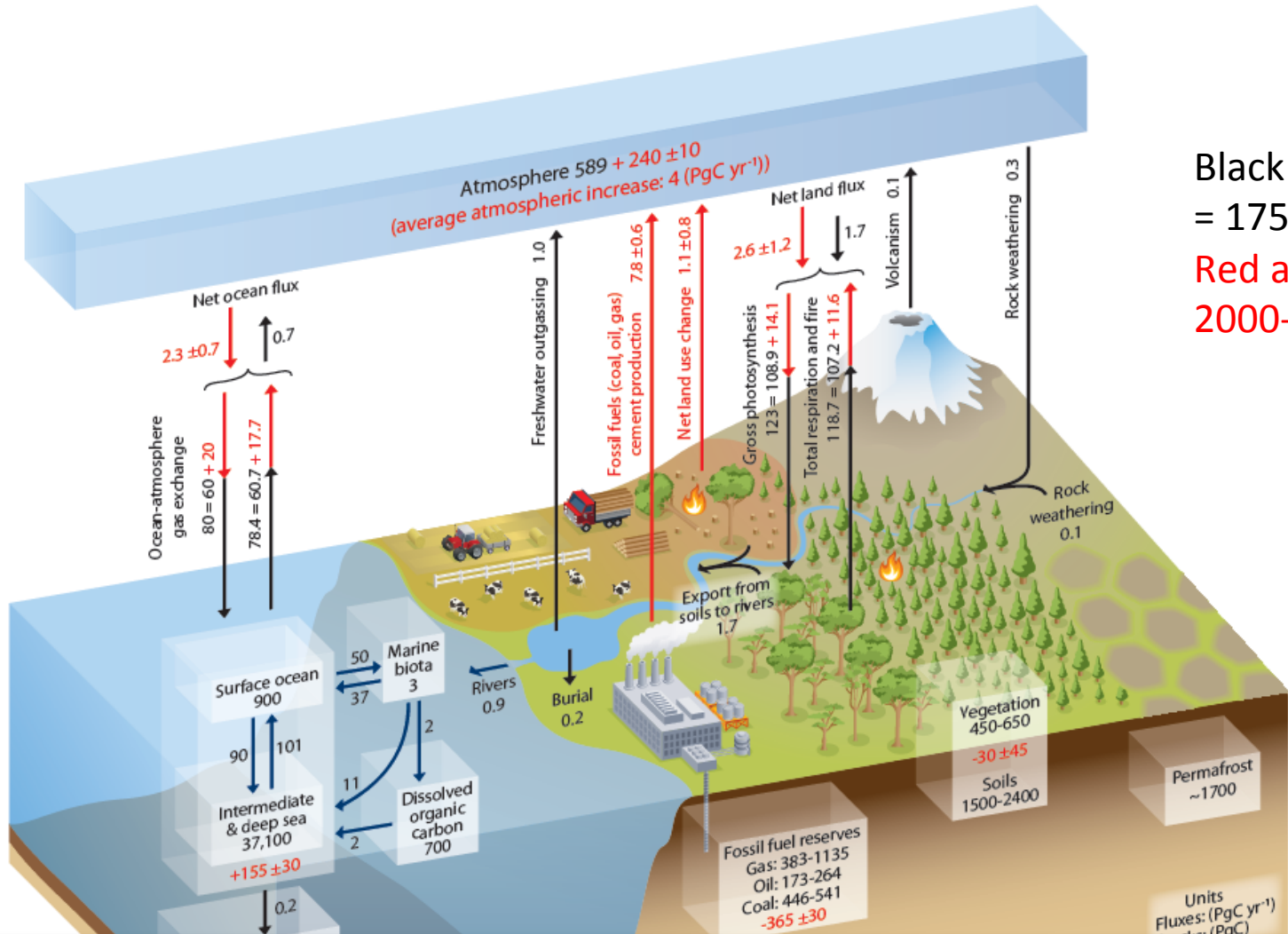


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$$d[\text{CO}_2]/dt = E_{\text{FF}} + E_{\text{LUC}} - S_{\text{O}} - S_{\text{L}}$$

Figure from Ciais et al. (2013)

Global carbon-cycle overview

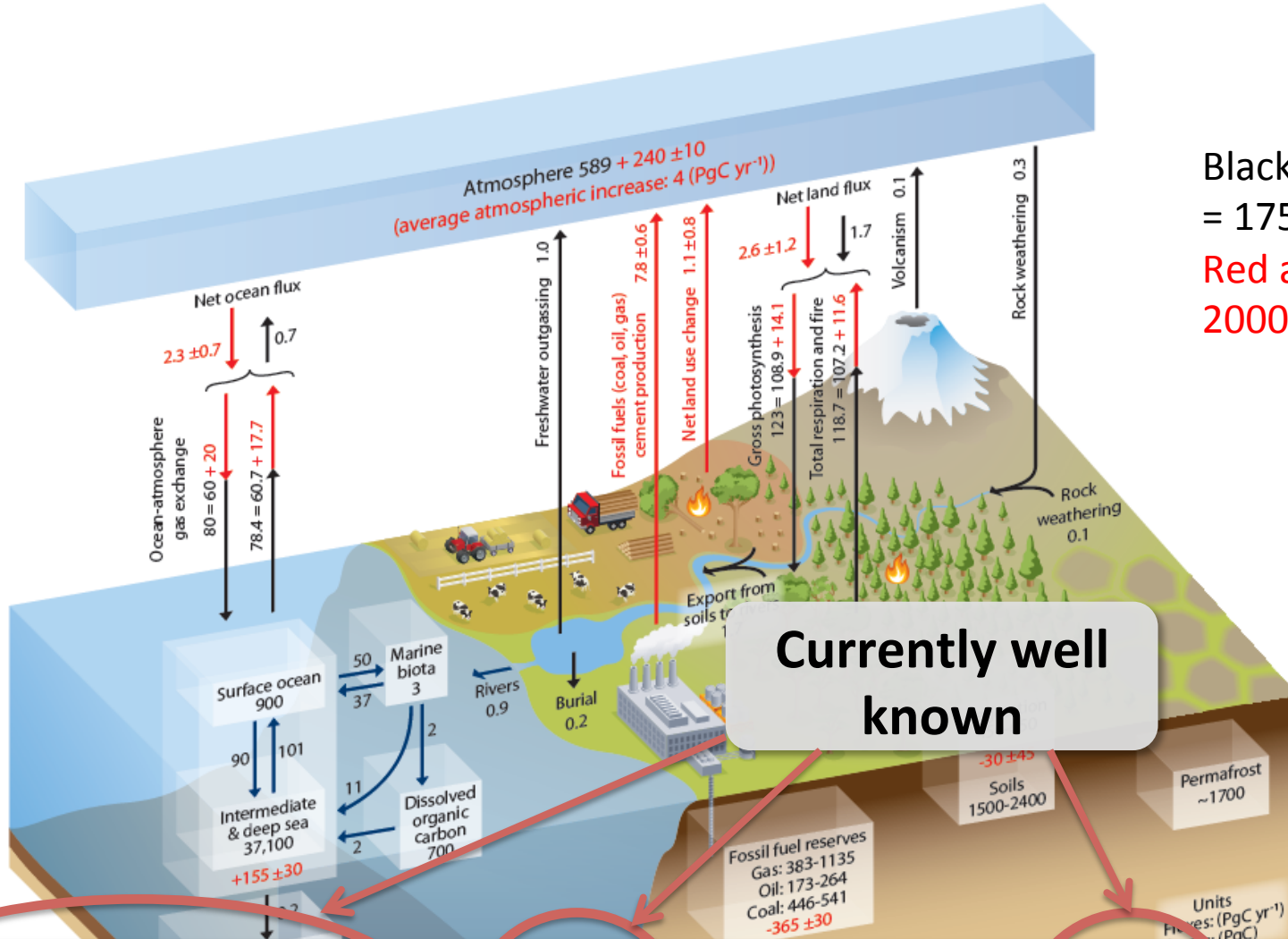


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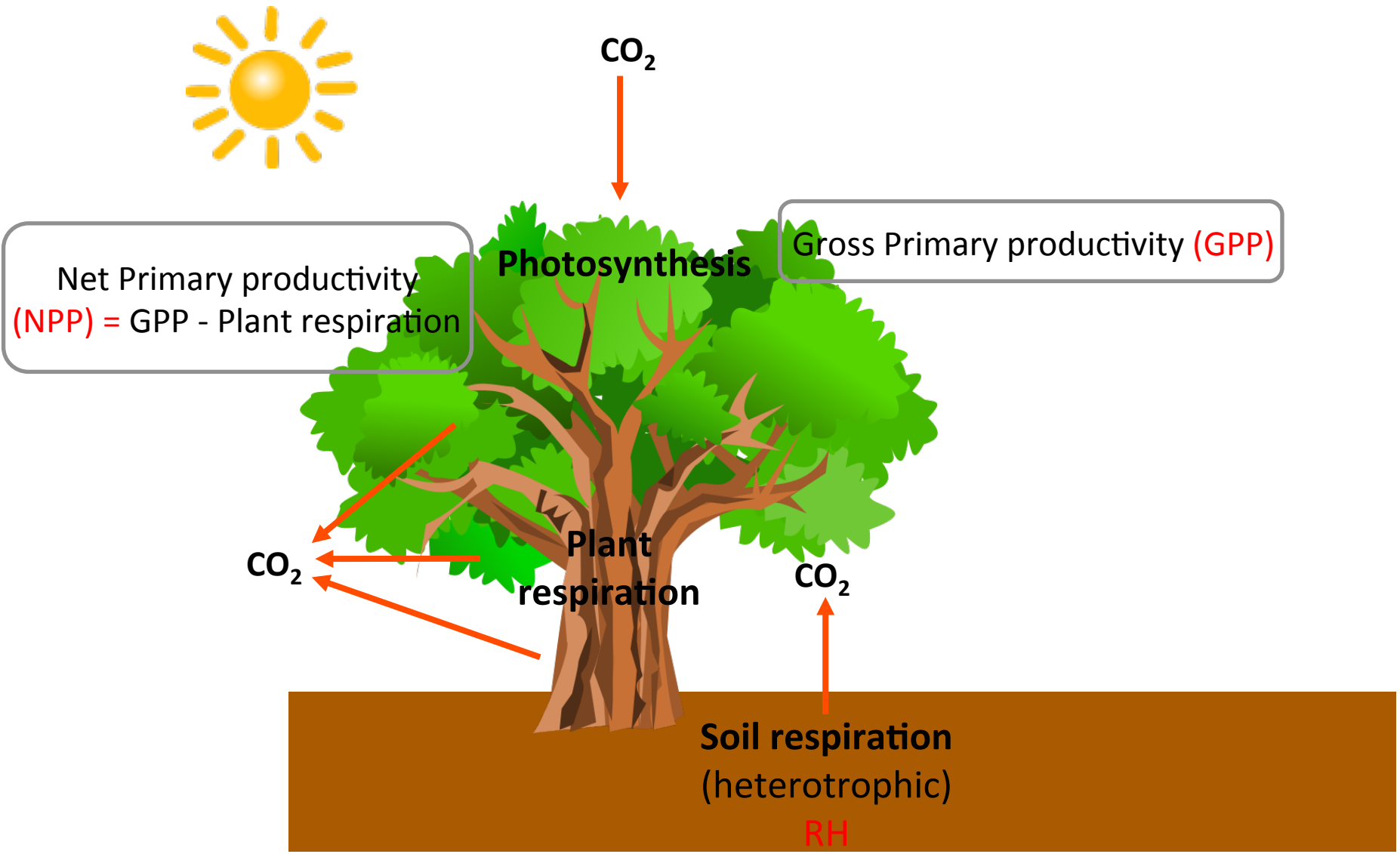
Currently well known

$$d[\text{CO}_2]/dt = E_{\text{FF}} + E_{\text{LUC}} - S_{\text{O}} - S_{\text{L}}$$

Figure from Ciais et al. (2013)

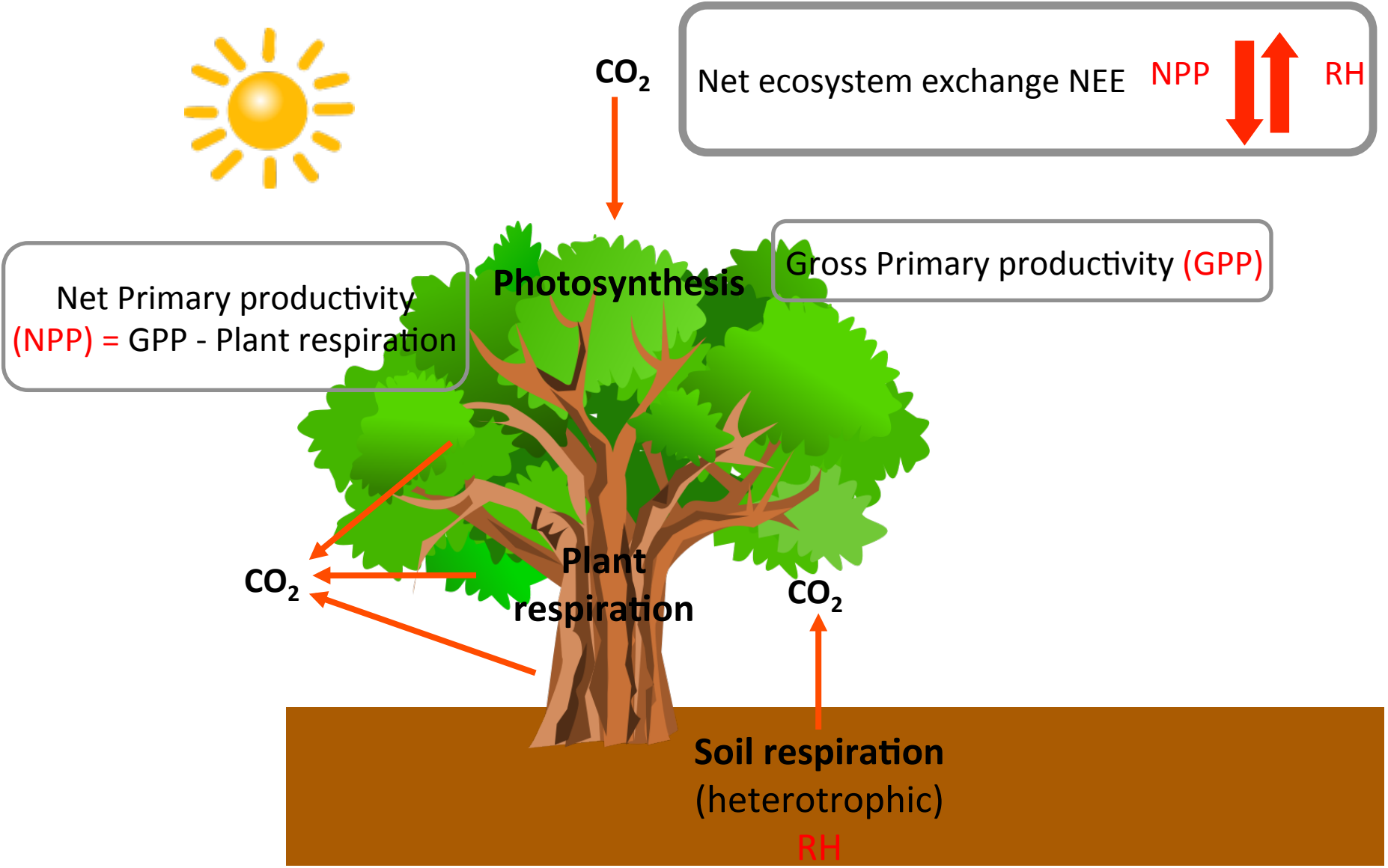
Global carbon-cycle overview

Basic elements of the terrestrial carbon cycle:



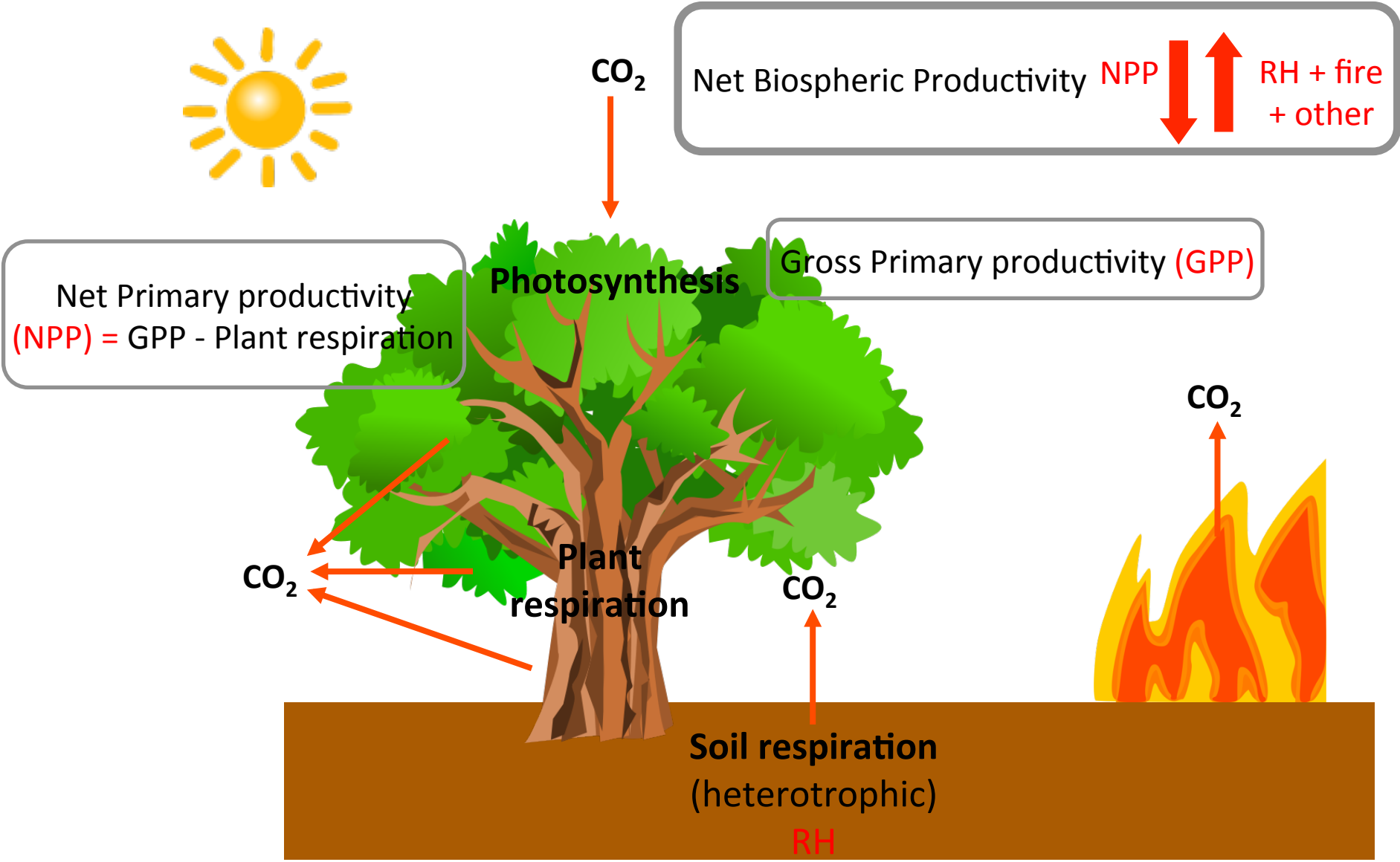
Global carbon-cycle overview

Basic elements of the terrestrial carbon cycle:



Global carbon-cycle overview

Basic elements of the terrestrial carbon cycle:



Global carbon-cycle modelling tools

Non-exhaustive overview of available measurements for terrestrial carbon cycle

Flux towers

- GPP
- NEE
- Total respiration

Spatial coverage: point

Frequency: <1 sec

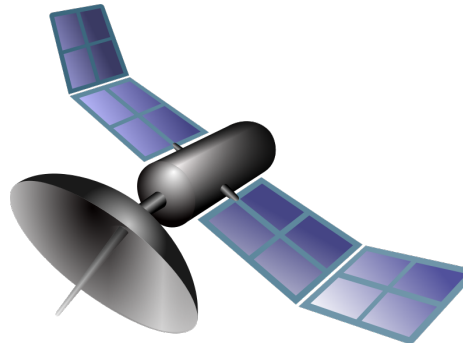


Satellites

- NDVI
- Canopy height
- GPP
(derived, e.g. from NDVI)

Spatial coverage: regional/
global

Frequency: days-weeks

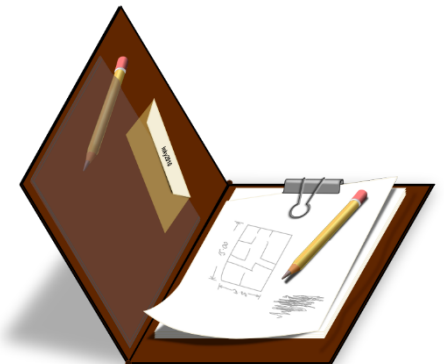


Inventories

- Biomass
- Growth rates
- Allometry

Spatial coverage: plot-
landscape

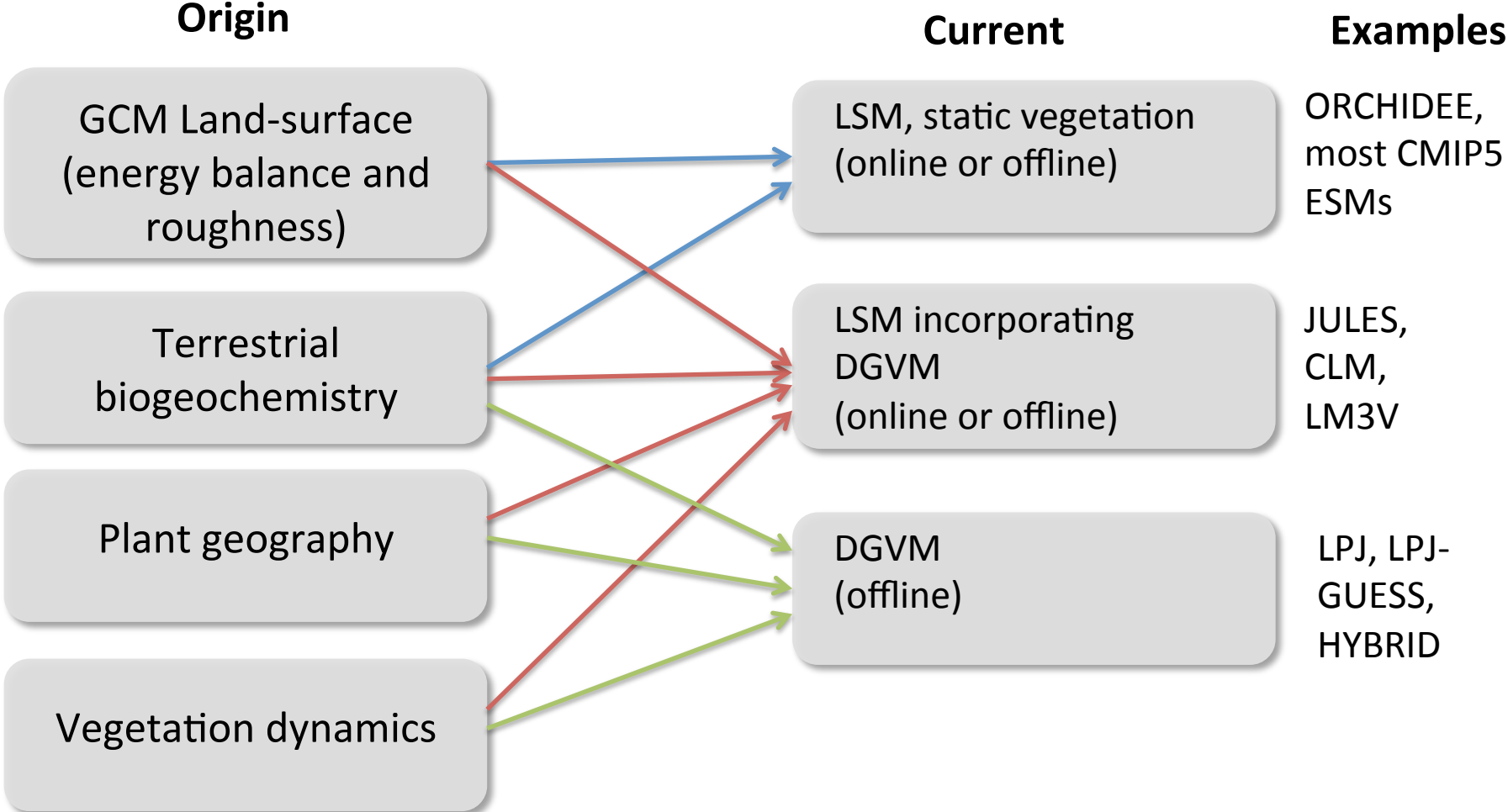
Frequency: years



Global carbon-cycle modelling tools

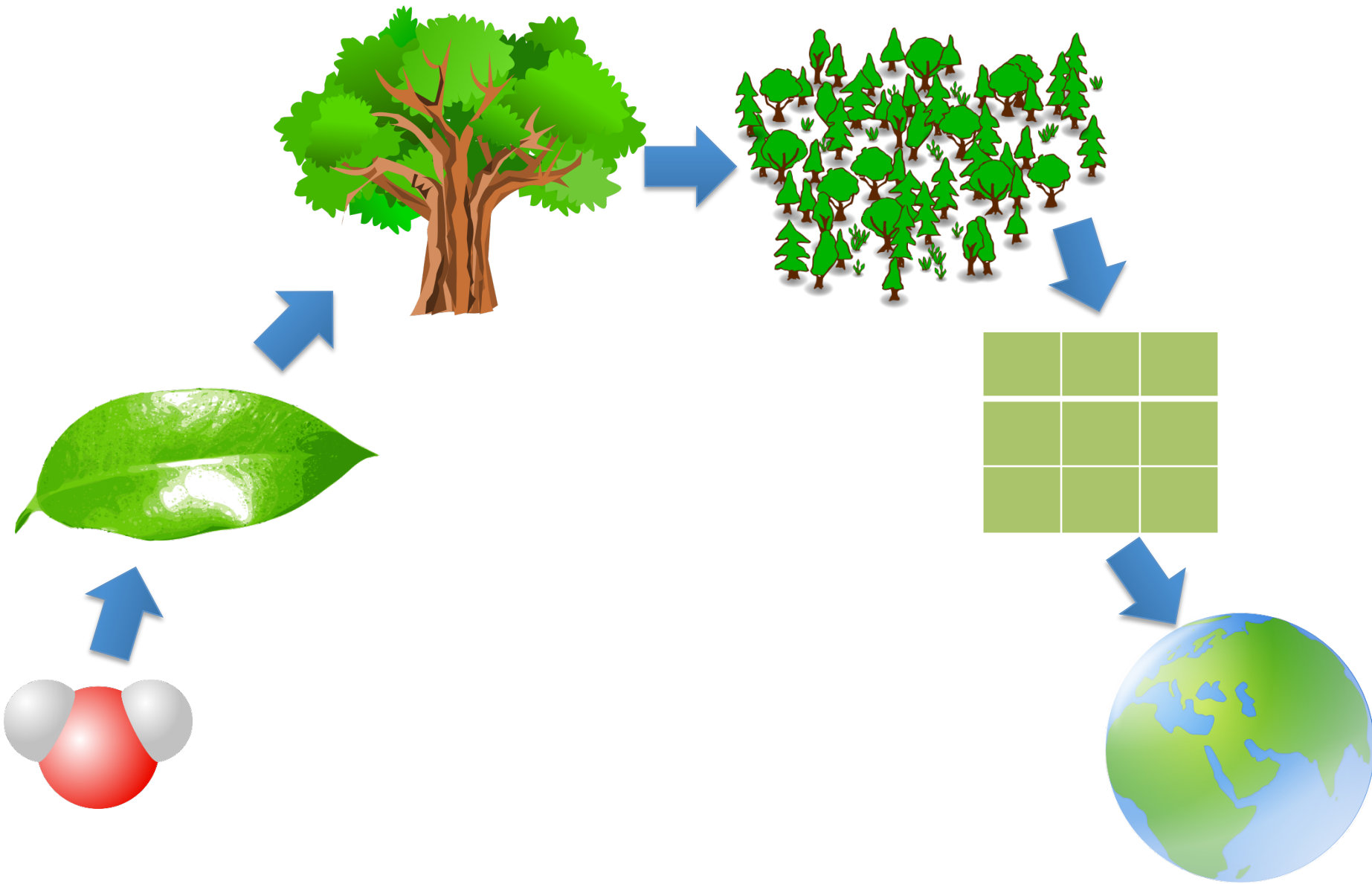
Typically we use process-based models to study the terrestrial biosphere.

The main models in use today have evolved from a variety of backgrounds, which make each differentially suited to the study of different aspects of the biosphere.



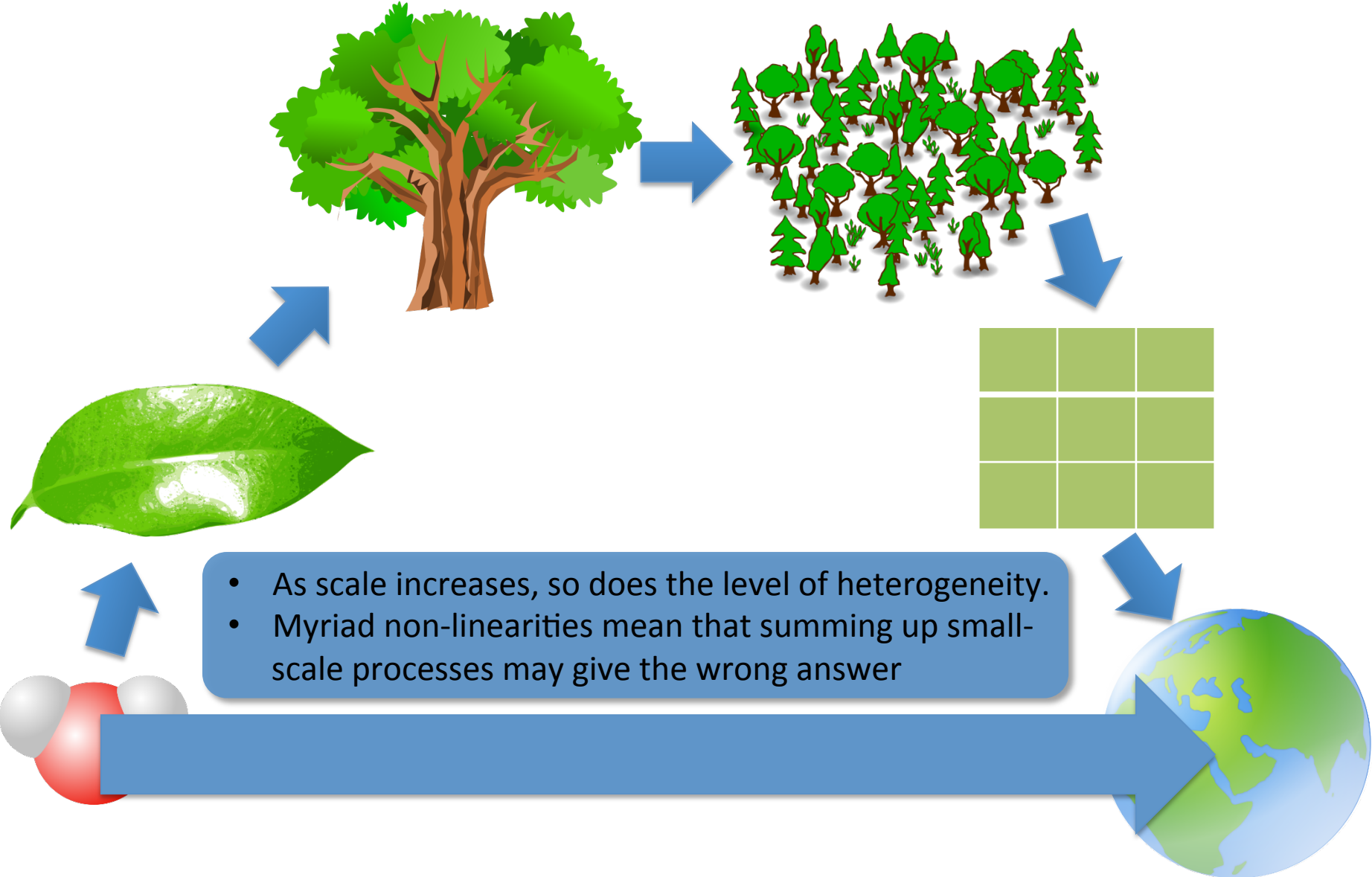
Global carbon-cycle modelling tools

Terrestrial biosphere models must span a range of spatial scales...



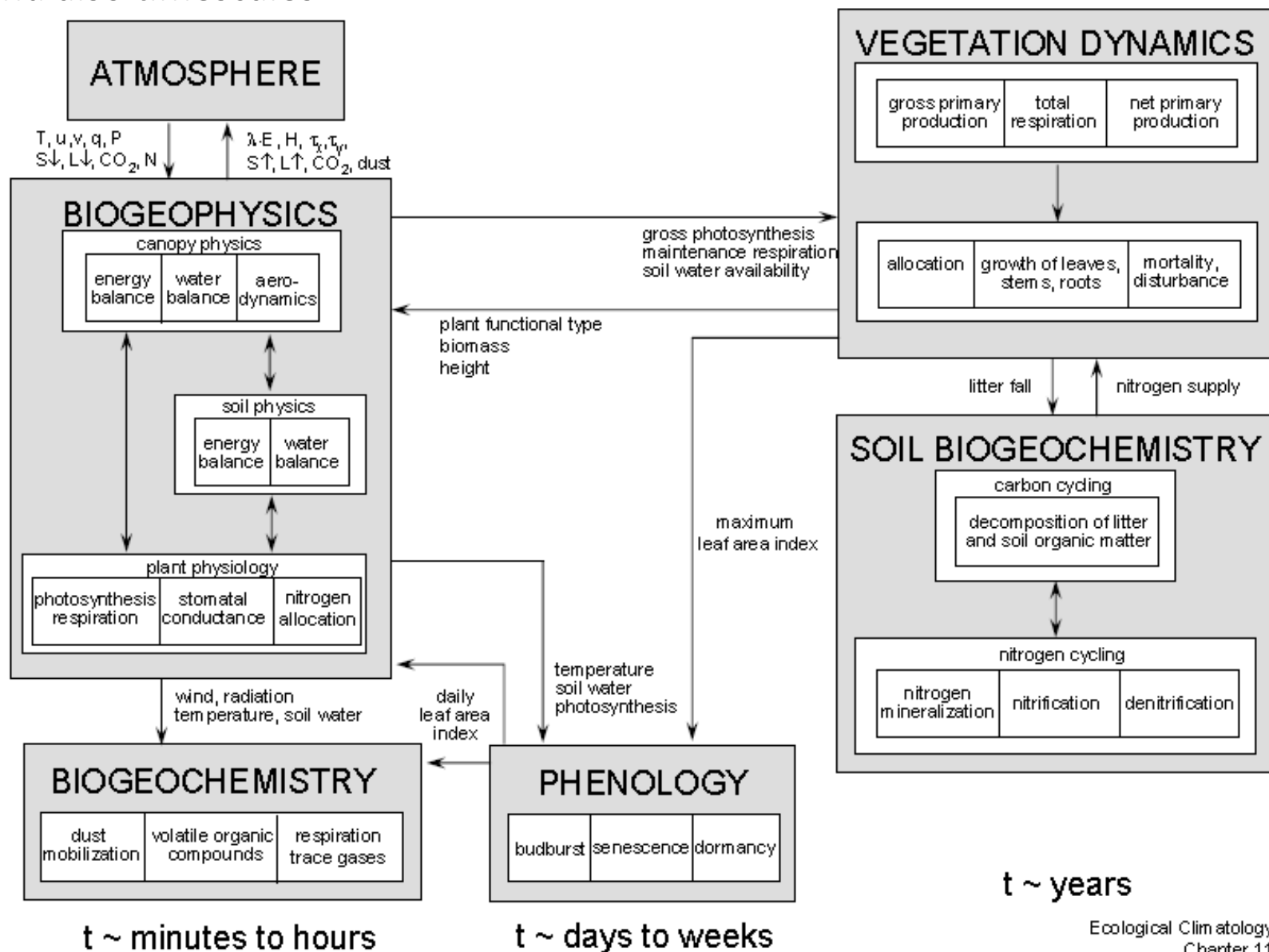
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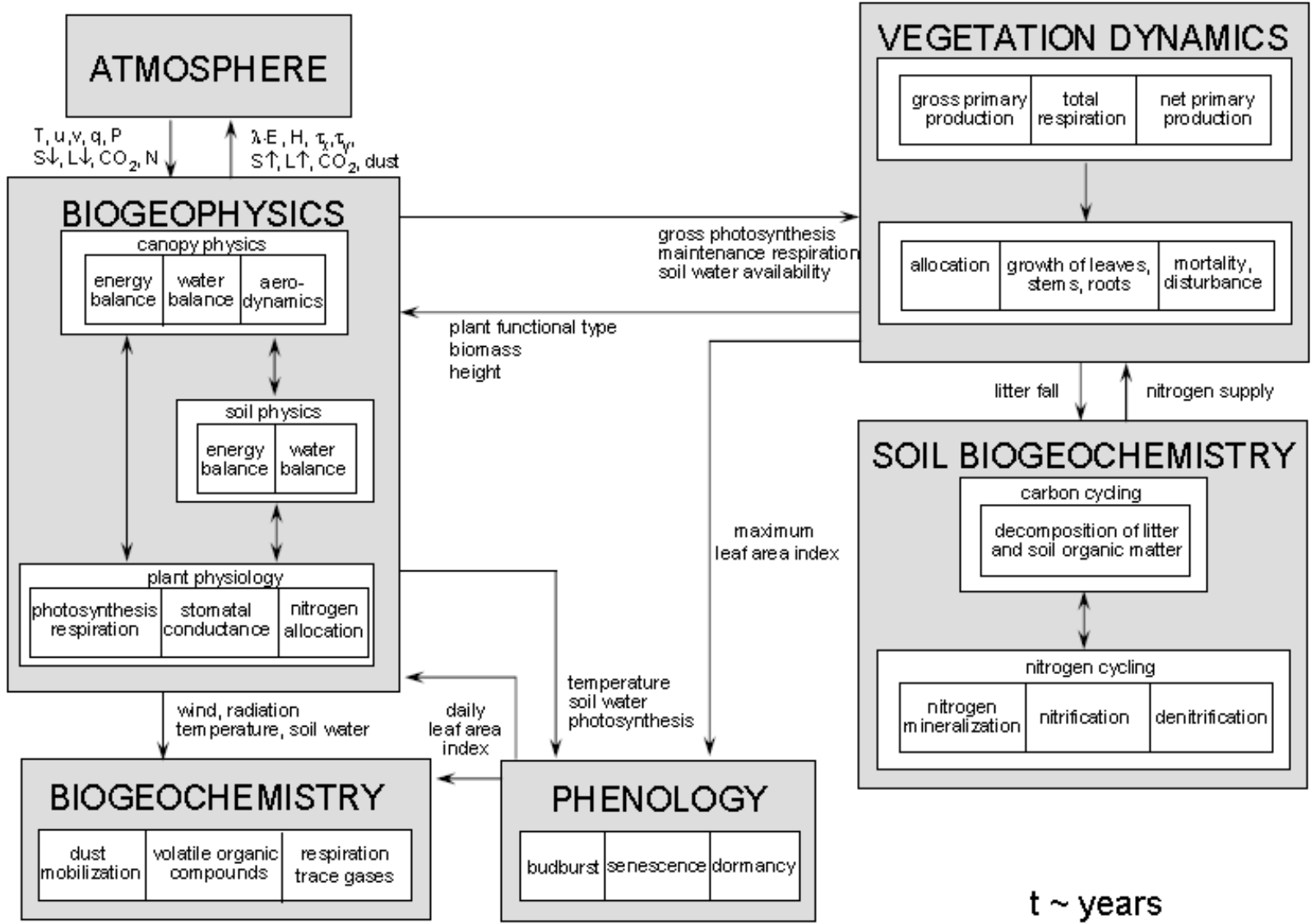
Global carbon-cycle modelling tools

... and also timescales



Global carbon-cycle modelling tools

... and also timescales



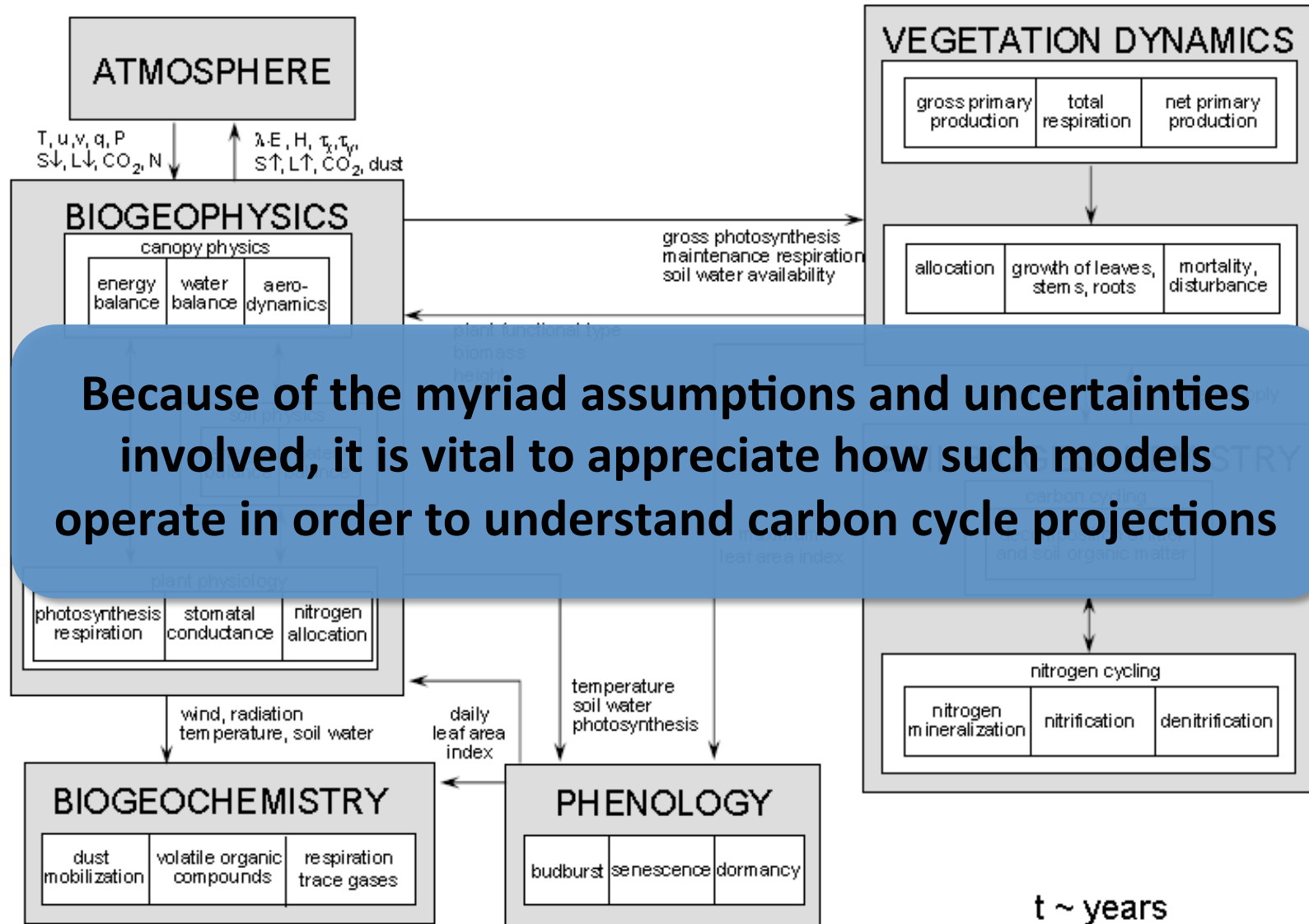
Often daily in DGVMs

t ~ days to weeks

t ~ years

Global carbon-cycle modelling tools

... and also timescales



Often daily in DGVMs

Overview of the remainder of this talk

Representations of key ecosystem processes in global terrestrial models

- Boundary conditions
- Primary production
- Respiration
- Structures and pools
- Species and vegetation dynamics
- Managed landcovers

Projecting the global carbon cycle - examples

Boundary conditions

The key boundary conditions for terrestrial biosphere models are usually atmospheric:

- Incoming short-wave radiation
- Surface/Air temperature (sometimes surface temp. is calculated explicitly from other variables)
- Precipitation
- Atmospheric CO₂ mixing ratio
- Incoming long-wave radiation
- Humidity
- Wind-speed
- Nitrogen/Phosphorus deposition

Others are:

- Soil parameters
- Vegetation type maps (not in DGVMs)
- Landcover/land-use/management variables

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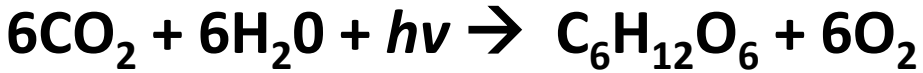
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Others are:

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Come from either historical reconstructions, or based on future scenarios (based on assumed human trajectories of e.g. emissions, management)

Primary production: Farquhar model



$\lambda = 400\text{-}700 \text{ nm}$

Photosynthesis is modelled as minimum of two limiting rates:

Carboxylation capacity (Rubisco-limited rate)

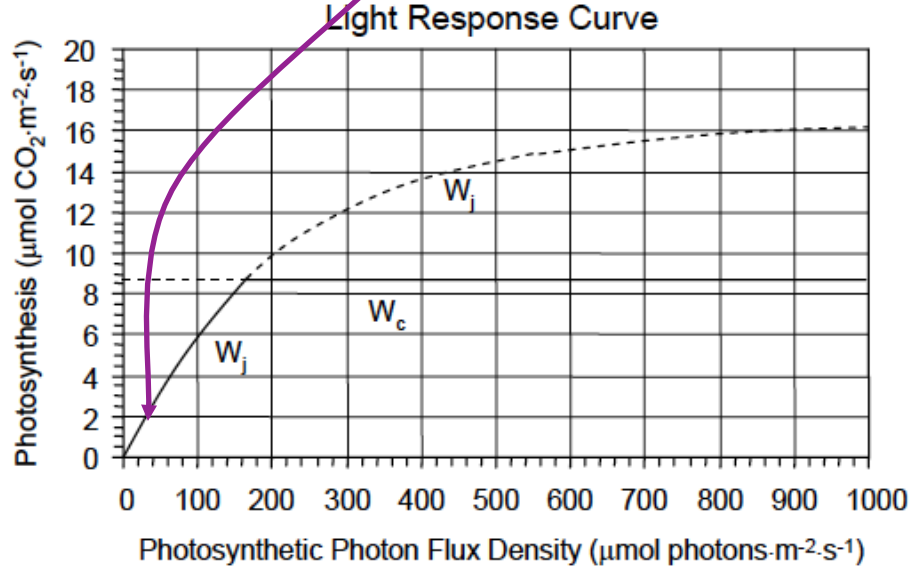
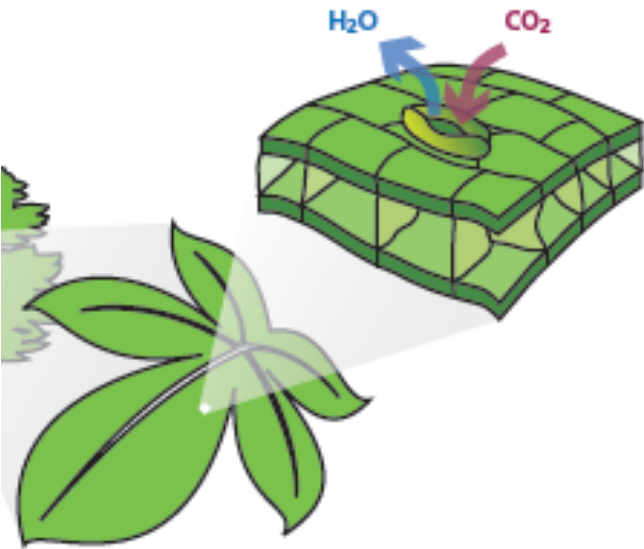
RuBP Regeneration (electron transport limited rate)

$$A = \min \left\{ \frac{V_{\max} (C_i - \Gamma^*)}{C_i + K_c (1 + O_i / K_o)}; \frac{J (C_i - \Gamma^*)}{4(C_i + 2\Gamma^*)} \right\}$$

Leaf-internal $[\text{CO}_2]$ $[\text{O}_2]$

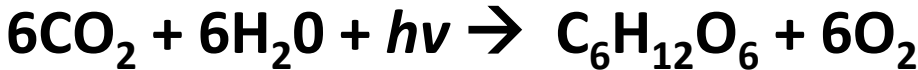
Constants

Photorespiration compensation point



Arneth et al., 2014, Slide A. Arneth

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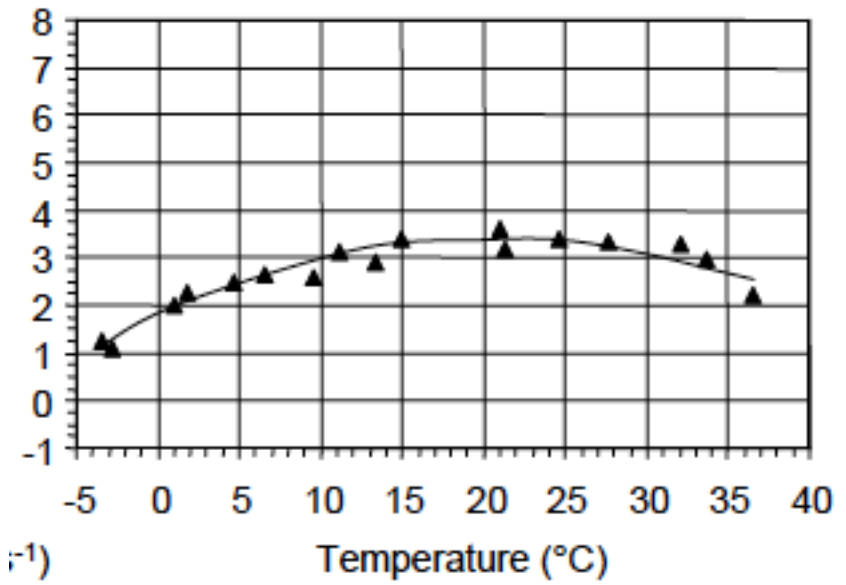
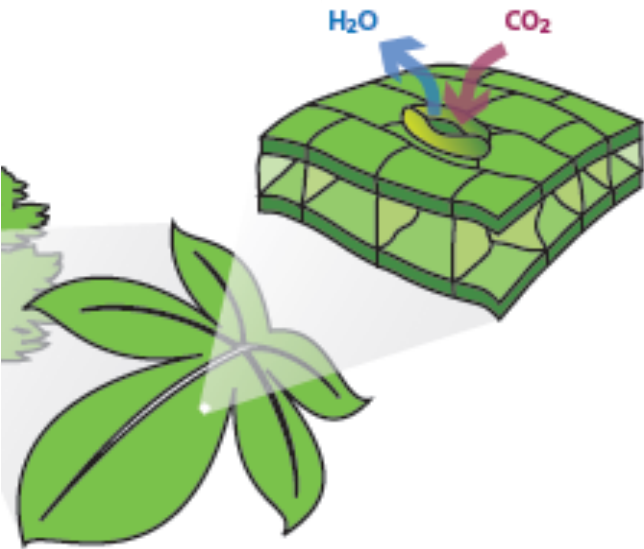
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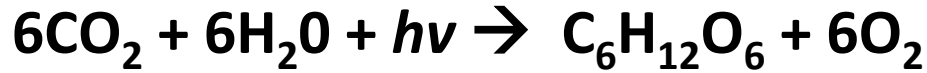
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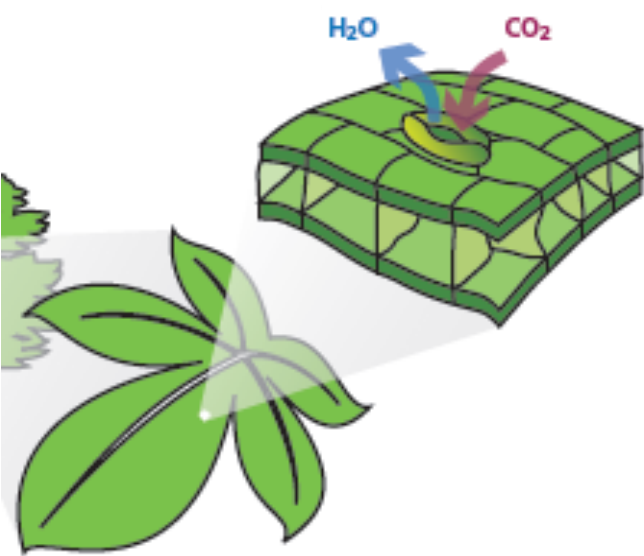
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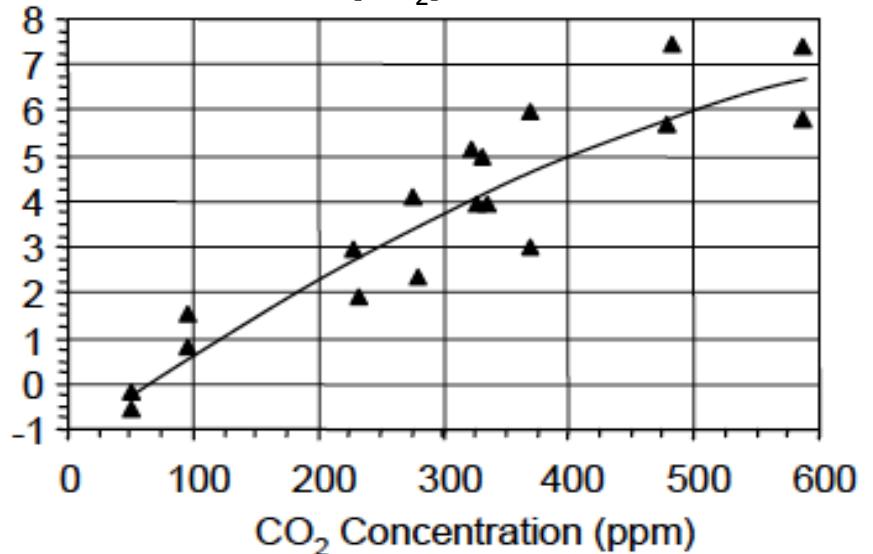
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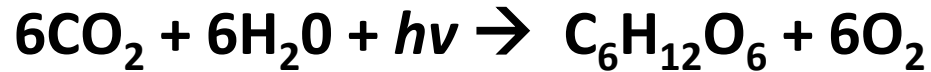


Net Photosynthesis ($\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$)

Sens. to $[\text{CO}_2]$ for Jack Pine



Primary production: Farquhar model



$\lambda = 400\text{-}700 \text{ nm}$

Photosynthesis is modelled as minimum of two limiting rates:

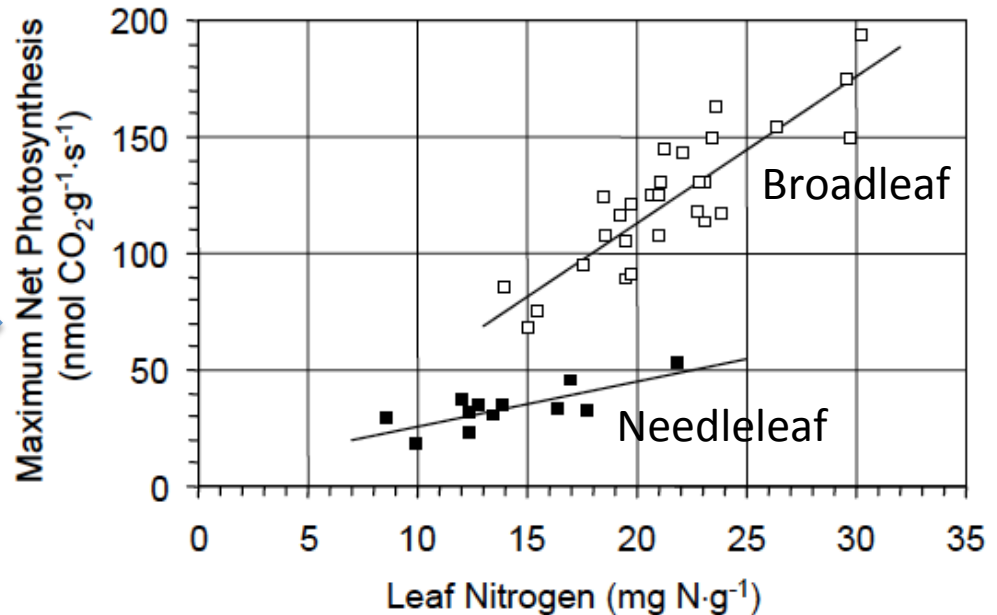
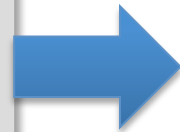
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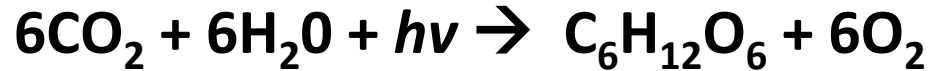
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Leaf-internal $[\text{CO}_2]$ $[\text{O}_2]$

Nitrogen is a fundamental control on photosynthesis. Key component of proteins, including Rubisco



Primary production: Farquhar model



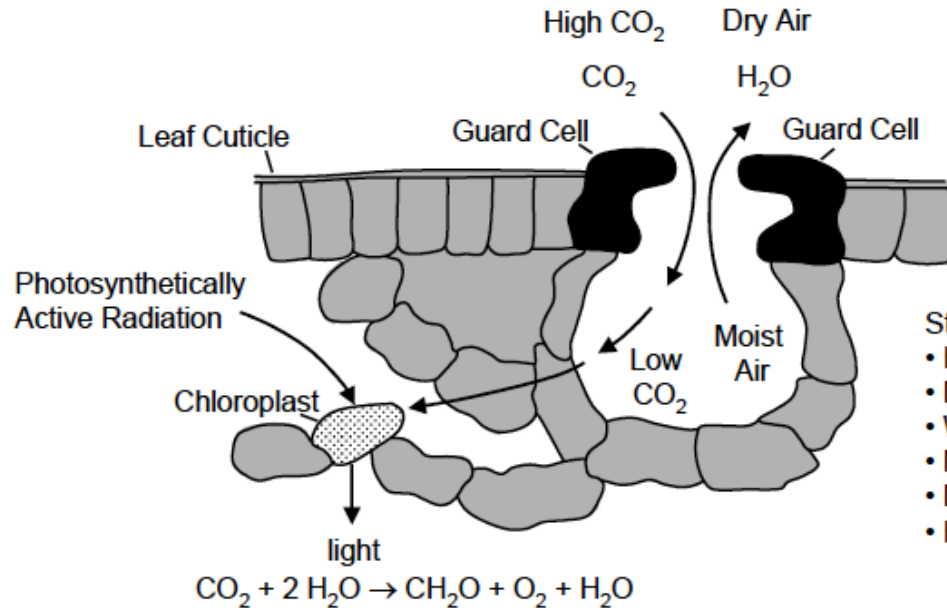
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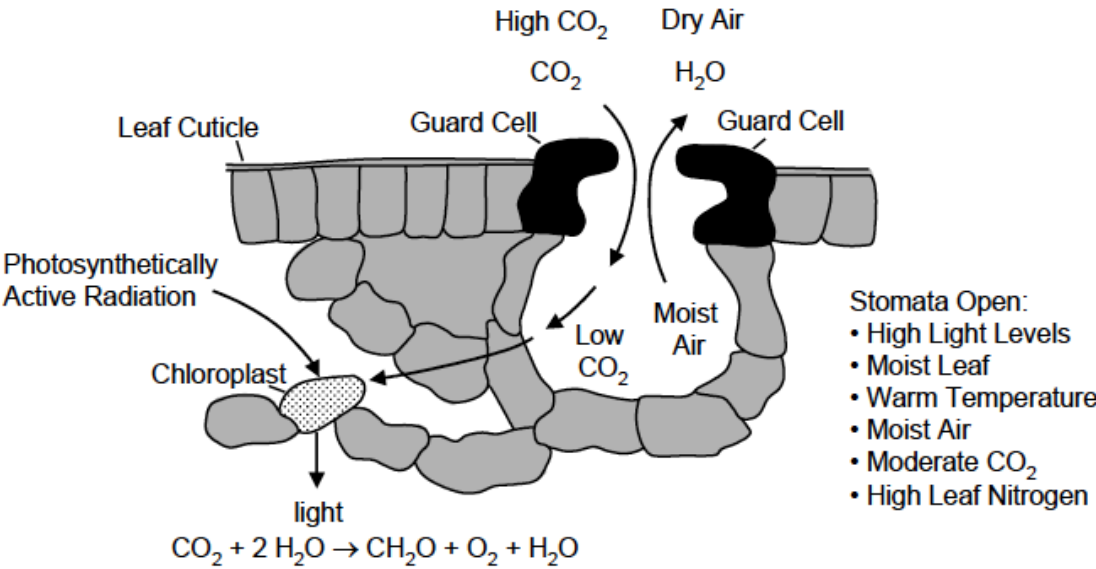
Amount of CO_2 available for photosynthesis is controlled by stomatal pores, but the trade-off is moisture loss



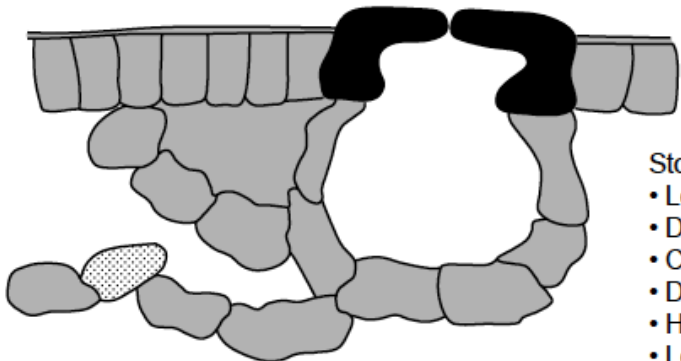
- Stomata Open:
- High Light Levels
 - Moist Leaf
 - Warm Temperature
 - Moist Air
 - Moderate CO_2
 - High Leaf Nitrogen

Primary production: Farquhar model

Stomatal Gas Exchange



- Stomata Open:**
- High Light Levels
 - Moist Leaf
 - Warm Temperature
 - Moist Air
 - Moderate CO₂
 - High Leaf Nitrogen



- Stomata Close (Smaller Pore Opening):**
- Low Light Levels
 - Dry Leaf
 - Cold Temperature
 - Dry Air
 - High CO₂
 - Low Leaf Nitrogen

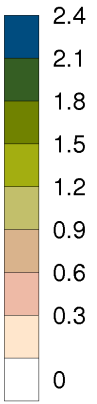
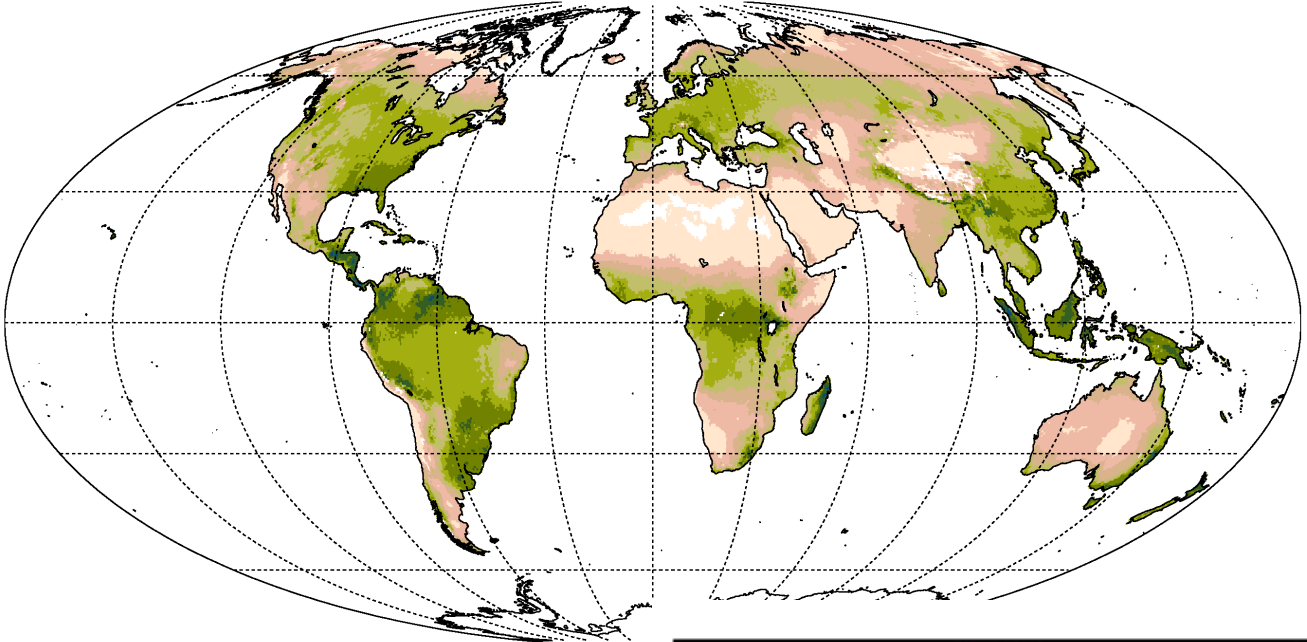
Optimality assumption:
Plants adjust stomatal conductance to optimise CO₂ intake vs water loss.

In LPJ:

$$g_c = g_{\min} + \frac{1.6A}{C_a (1-\alpha)}$$

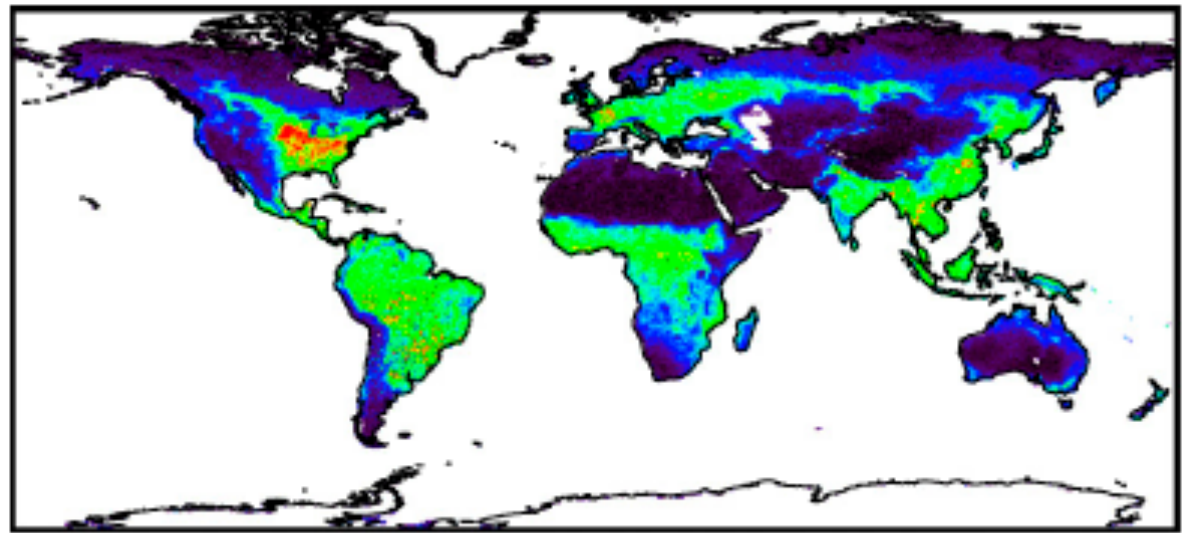
α is reduced under water stress

Primary production



GPP
from LPJ-GUESS
1981-2000
(kg C m⁻² a⁻¹)

Chlorophyll fluorescence –
can be detected by
satellite. Indicator for
GPP



max(SIF) (mW/m²/sr/nm)
Guanter et al. (2014)

A horizontal color scale legend for max(SIF). The scale ranges from 0.0 (dark purple) to 4.5 (red). Intermediate values are marked at 0.9, 1.8, 2.7, and 3.6.

Autotrophic respiration

Plant respiration usually divided into growth and maintenance respiration

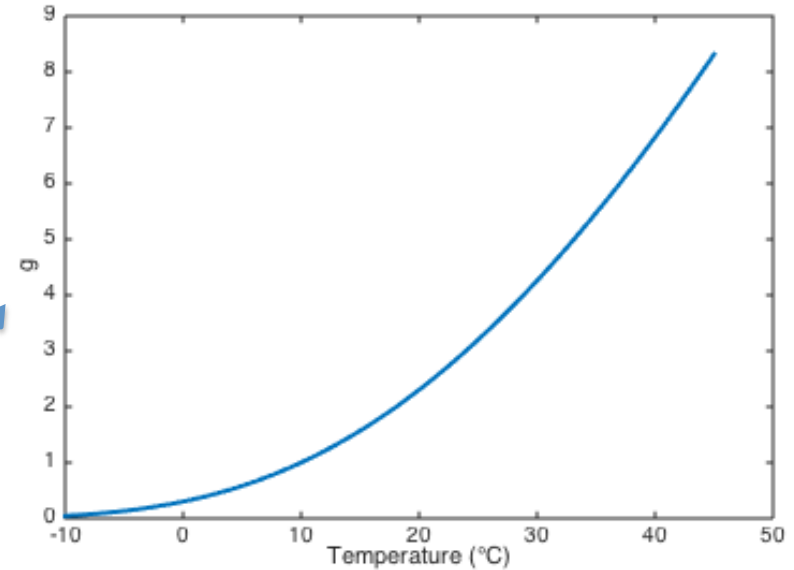
Example functions from LPJ/LPJ-GUESS:

$$R_{\text{leaf}} = r \cdot \frac{C_{\text{leaf}}}{\text{cn}_{\text{leaf}}} \phi \cdot g(T)$$

$$R_{\text{sapwood}} = r \cdot \frac{C_{\text{sapwood}}}{\text{cn}_{\text{sapwood}}} g(T)$$

$$R_{\text{root}} = r \cdot \frac{C_{\text{root}}}{\text{cn}_{\text{fineroot}}} \phi \cdot g(T_{\text{soil}})$$

Rate is tissue dependent – related to nitrogen content



$$g(T) = \exp \left[308.56 \cdot \left(\frac{1}{56.02} - \frac{1}{(T + 46.02)} \right) \right]$$

Strong dependence on temperature

Growth:

$$R_g = 0.25 \times (GPP - R_{\text{leaf}} + R_{\text{sapwood}} + R_{\text{root}})$$

Further reading: Thornley and Cannell (2000), Annals of Botany

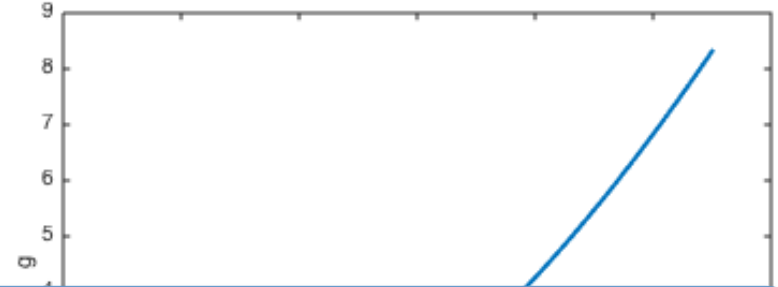
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BUT: High simplification of underlying biochemistry. Acclimatisation could moderate response in warm future climates. Model appears to fail under drought conditions.

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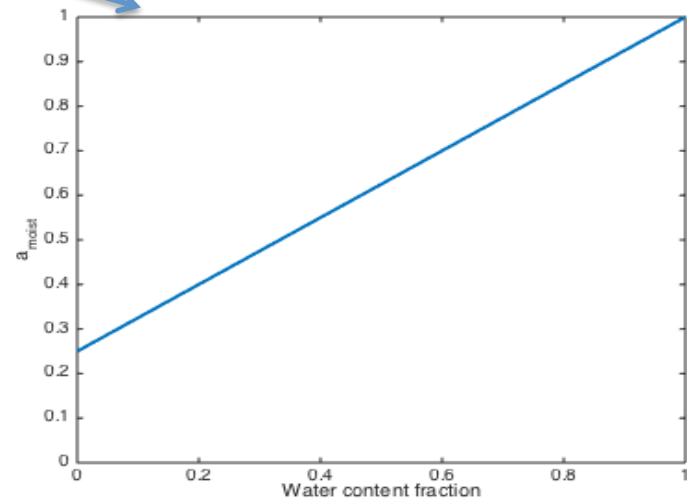
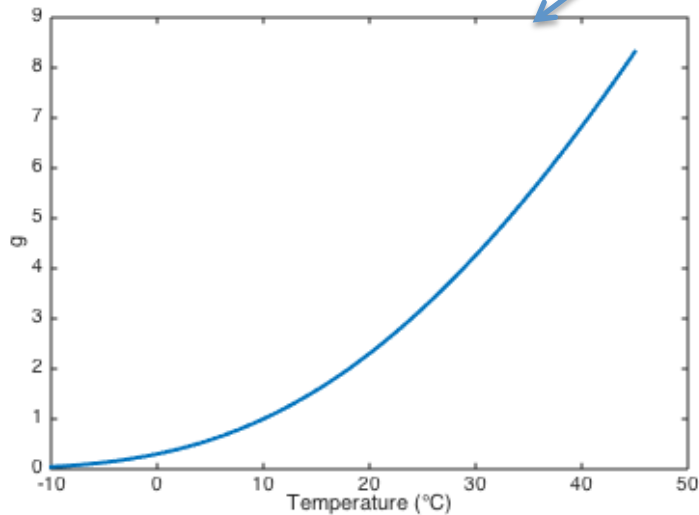
Heterotrophic respiration

Respiration by decomposers in the soil.

Typically global models use simple lifetime functions modified by temperature and moisture:

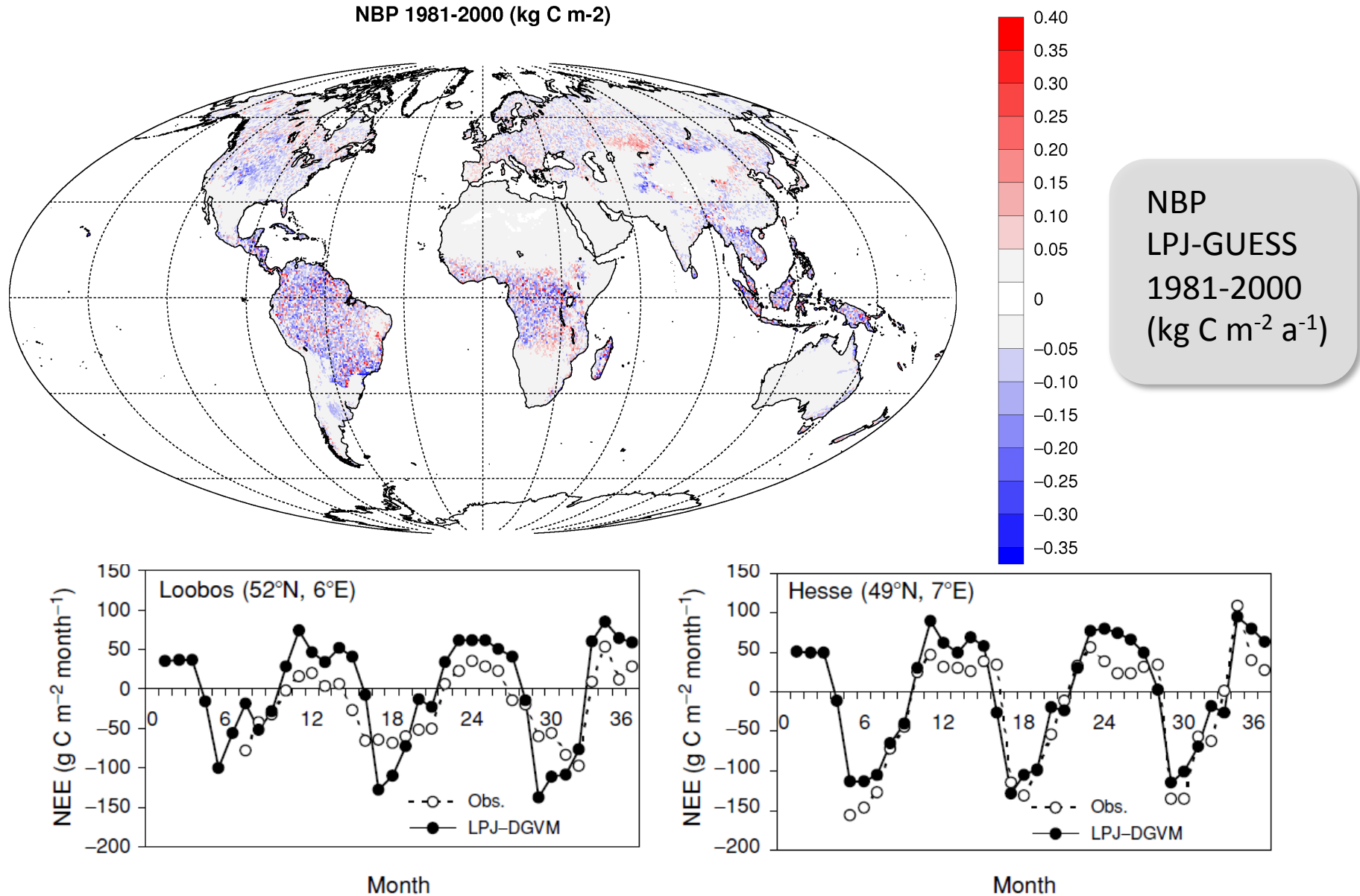
$$R_H = \exp(1 / \tau) \cdot g_T \cdot g_{moist}$$

$$\tau \approx 2-1000 \text{ years}$$



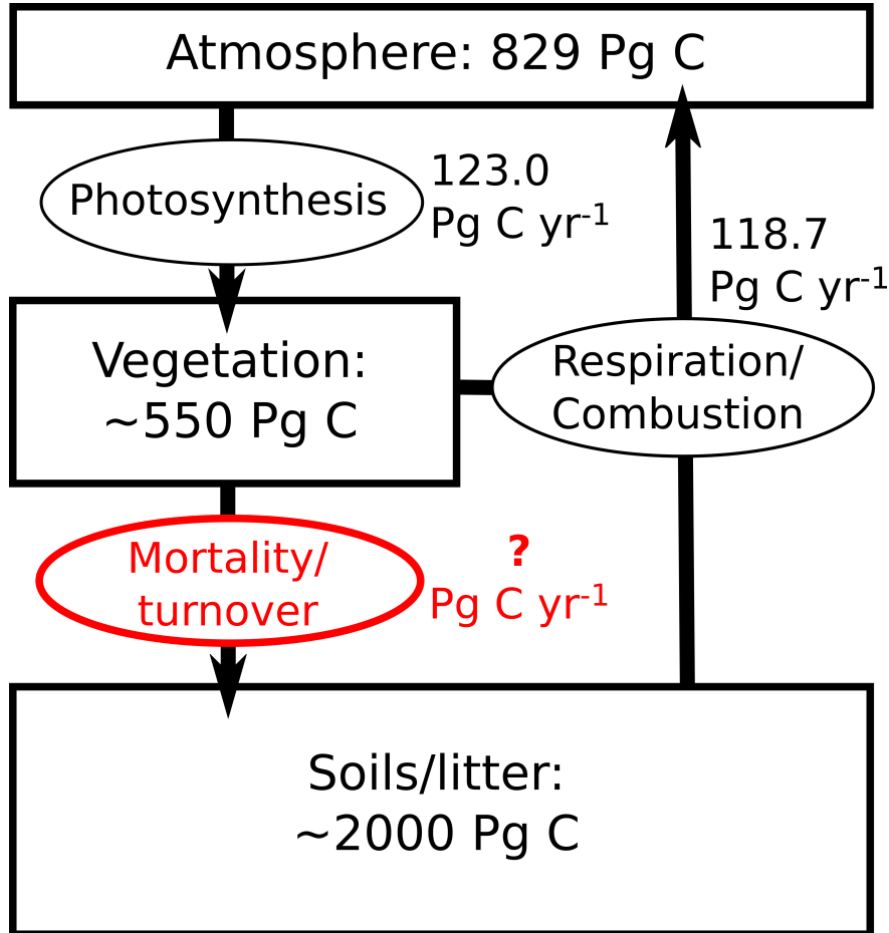
Heavily over-simplified in most global models. Is actually a function of a whole range of factors – bacteria, substrate, moisture, N, temperature. See, e.g. Koven et al. (2013)

Net ecosystem flux balance



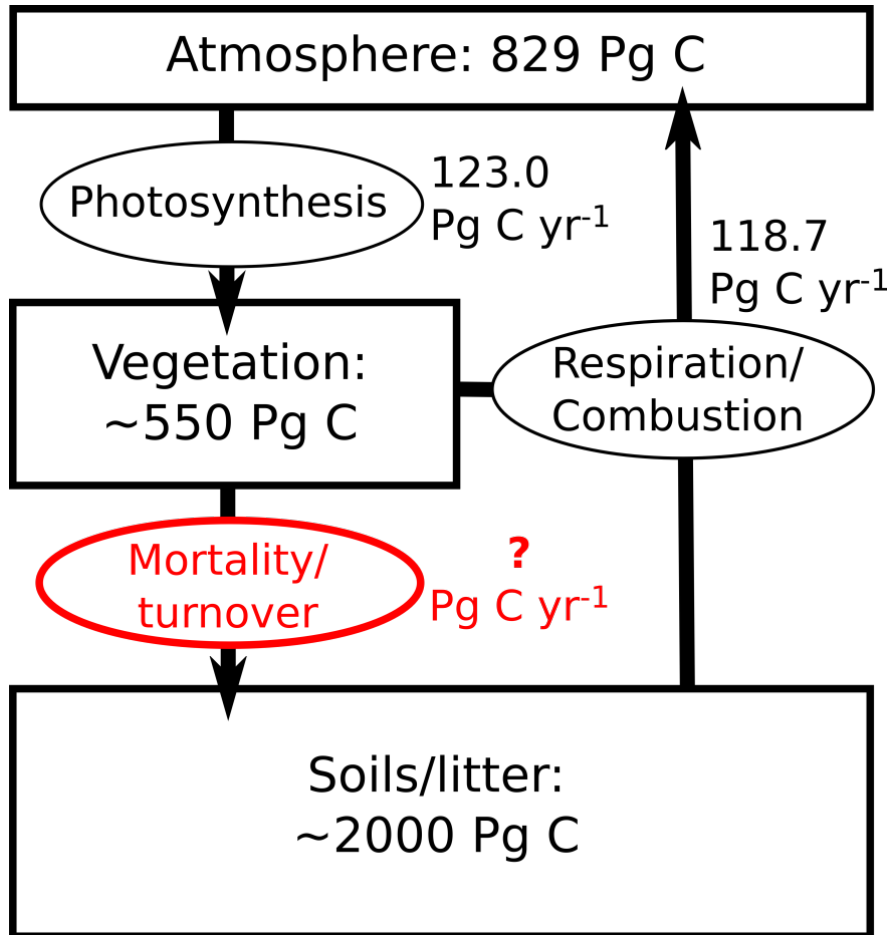
Model structure

We model the effect of these fluxes on terrestrial carbon storage using the concept of pools.



Model structure

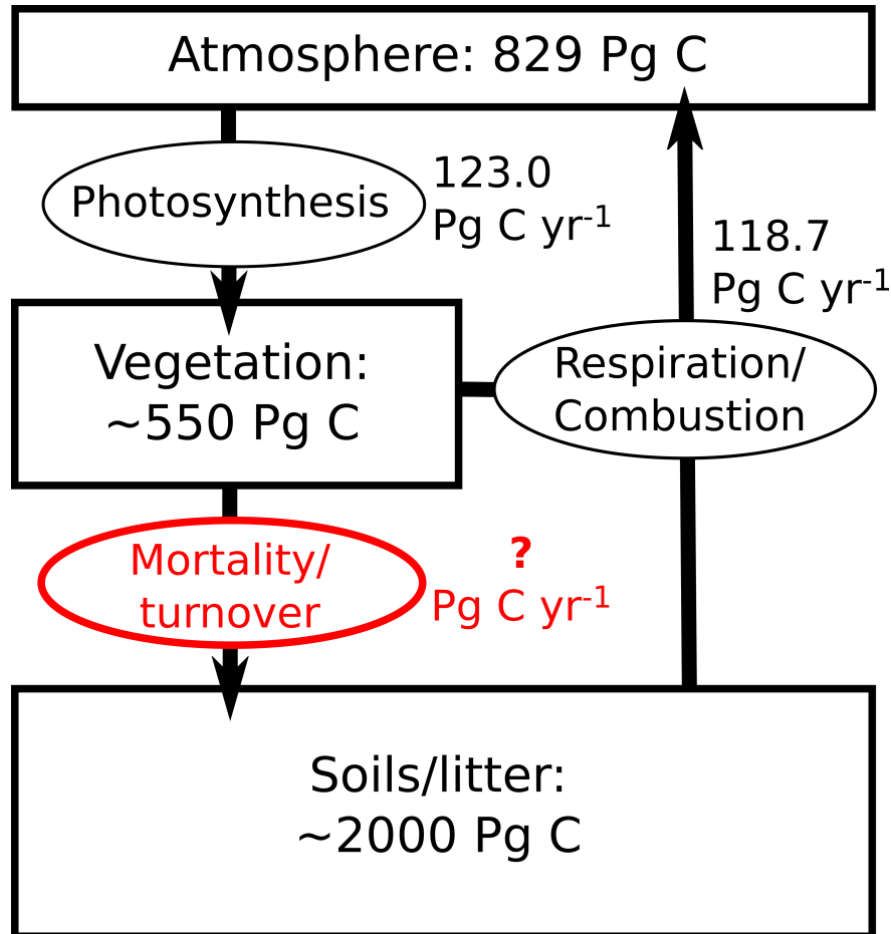
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Soils and litter are often split into several different sub-pools depending on their basic decomposition rates. Typical pool turnover times are a few years for litter, and 10-1000 years for soil pools

Model structure

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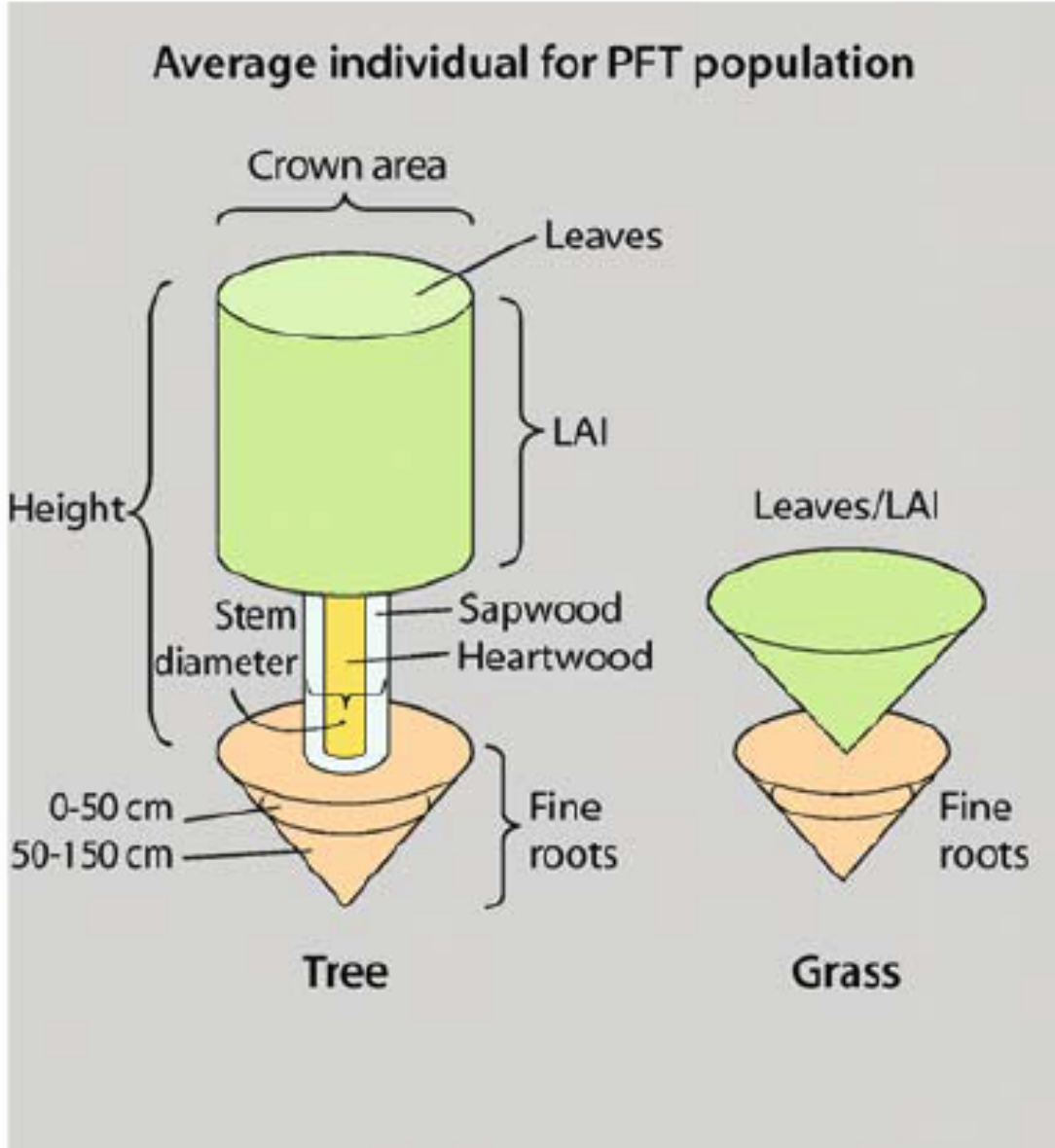


Vegetation pool may be one average individual, or a range of different individuals of different age and species, depending on the model



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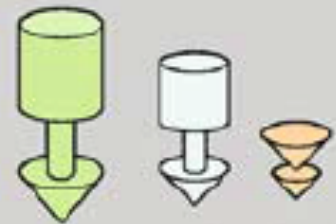
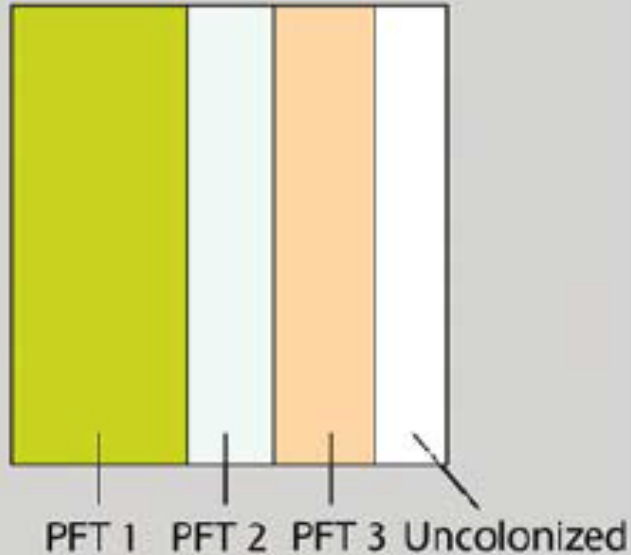
Model structure and vegetation dynamics



Model structure and vegetation dynamics

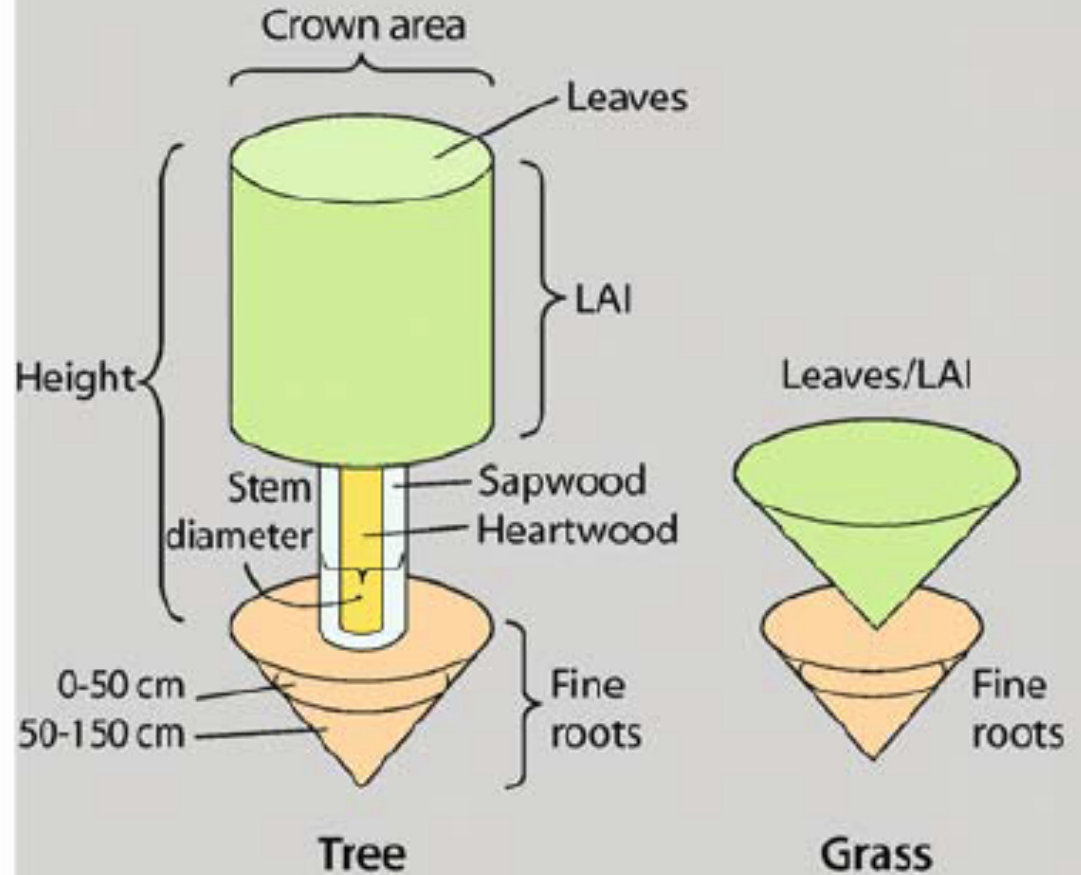
Modeled area (grid cell)
ca. 100-2500 km²

Fractional cover (FPC) x PFT

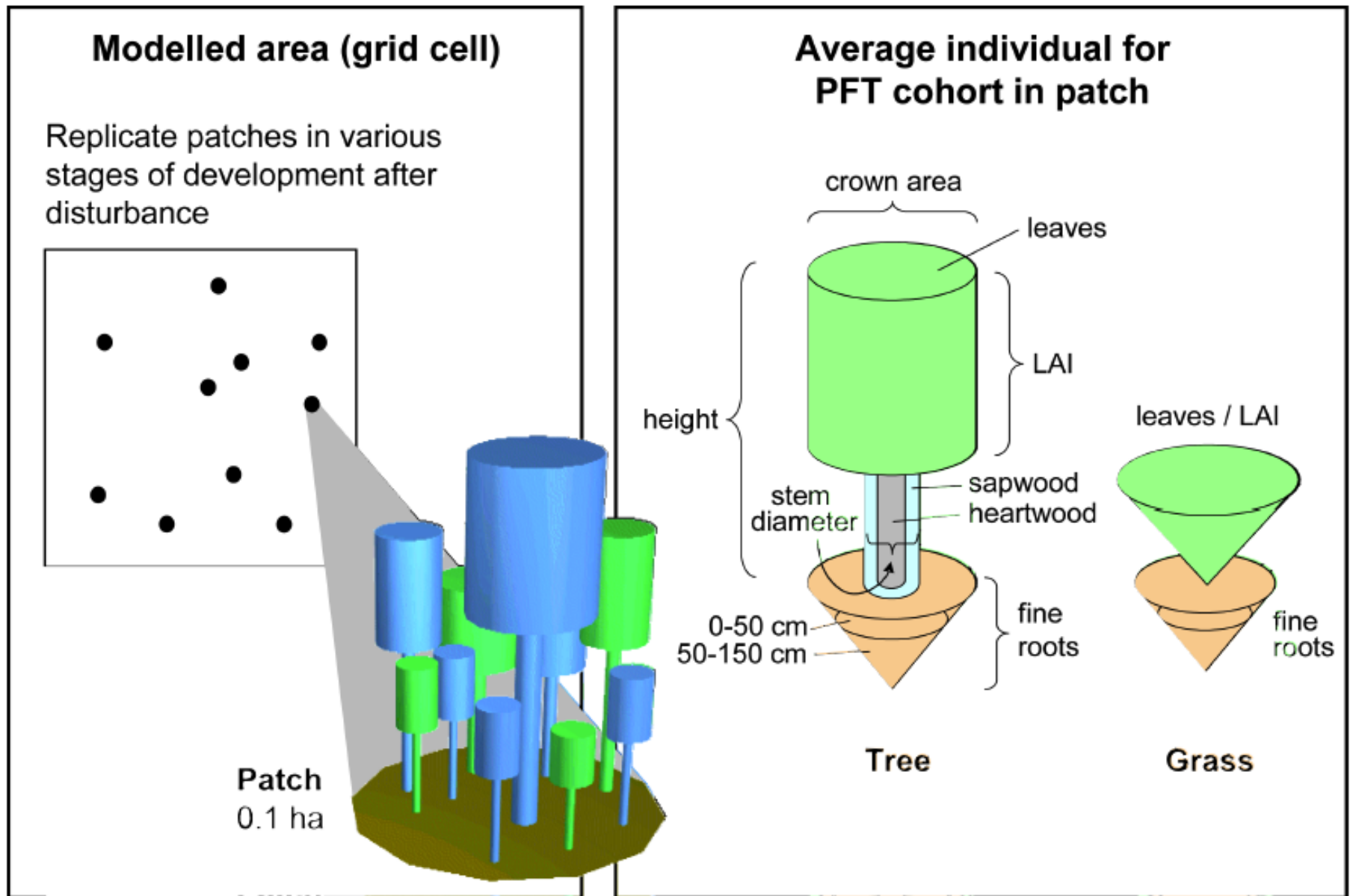


Average individual
for PFT population

Average individual for PFT population



Model structure and vegetation dynamics



Model structure and vegetation dynamics

Plants can also die. Typical death mechanisms related to resource availability or physiological limits include:

- Bioclimatic limits
- Negative productivity
- Growth efficiency threshold (biomass increase per unit leaf area)
- Maximum age
- Background rate
- Shading/competition (mortality increases with canopy cover)

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Do not reflect the actual mechanisms which lead to plant death

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Plants can also die due to ecosystem disturbances, e.g.

- Fire
- Wind-throw
- Insect attack

This is a major uncertainty in the understanding of ecosystem response to future environments. A doubling of mortality rates leads to a large drop in vegetation carbon stocks.

Model structure and vegetation dynamics

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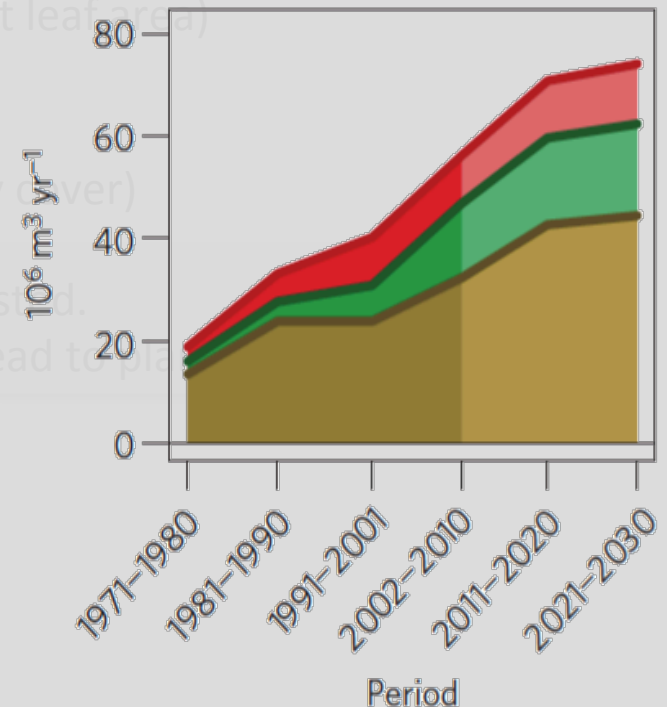
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Disturbance agent

- Forest fire
- Bark beetles
- Wind

Seidl et al., 2014

Model structure and vegetation dynamics

Plants can also die. Typical death mechanisms related to resource availability or physiological limits include:

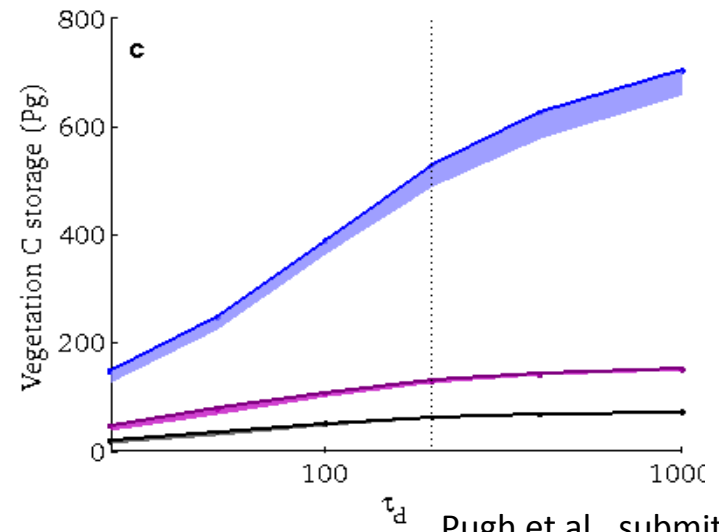
- Bioclimatic limits
- Negative productivity
- Growth efficiency threshold (biomass increase per unit leaf area)
- Maximum age
- Background rate
- Shading/competition (mortality increases with canopy cover)

These are logical, but not well tested.
Do not reflect the actual mechanisms which lead to plant death

Plants can also die due to ecosystem disturbances, e.g.

- Fire
- Wind-throw
- Insect attack

This is a major uncertainty in the understanding of ecosystem response to future environments. A doubling of mortality rates leads to a large drop in vegetation carbon stocks.

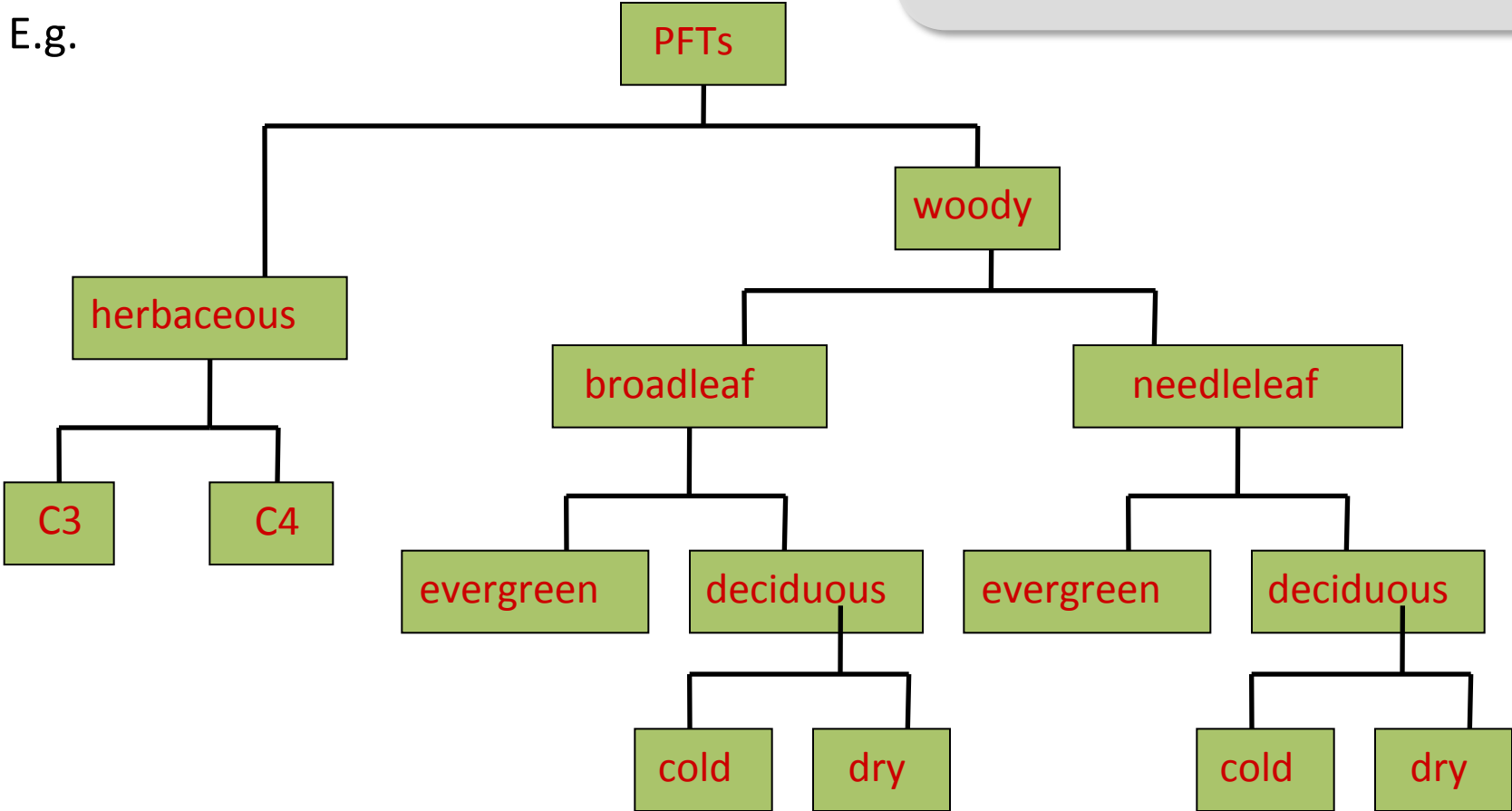


Species composition

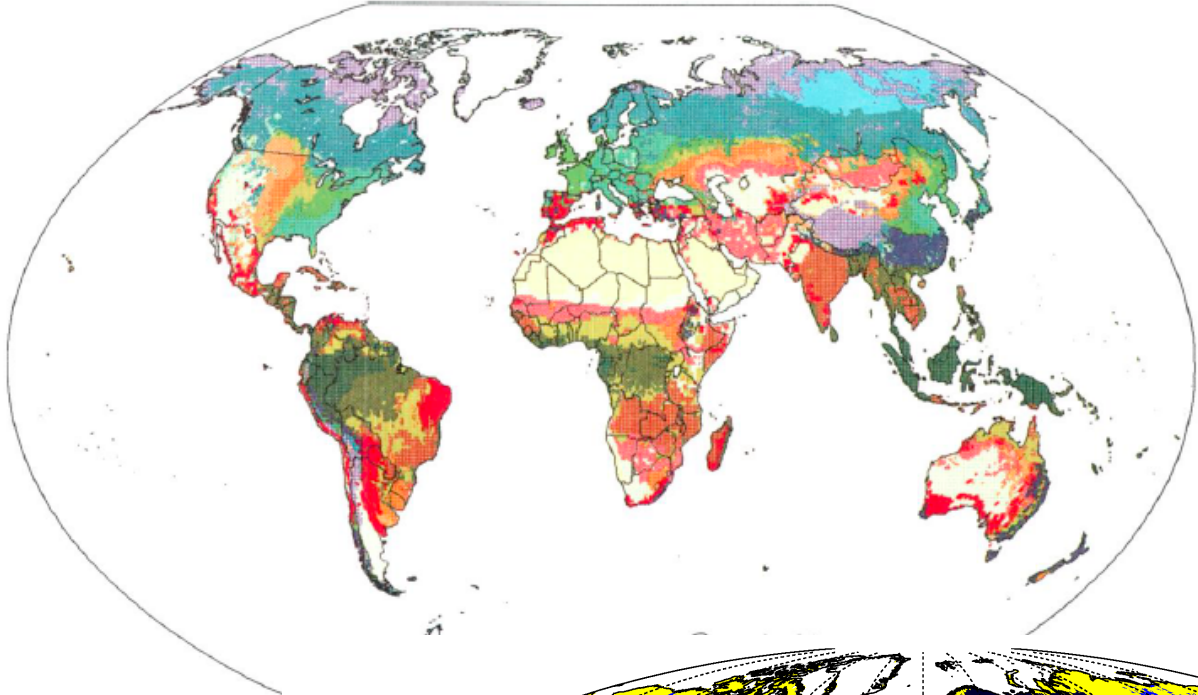
Huge range of species in reality. We have neither computational capacity or data to capture these in large-scale models. Typical approach is to classify species into PFTs.

E.g.

Distinguished by, e.g.
Plant physiology (C3/C4 photosynthesis)
Phenology (Evergreen/Deciduous)
Physiognomy (Woody/Herbaceous)
Bioclimatic limits (Cold/Heat tolerance)



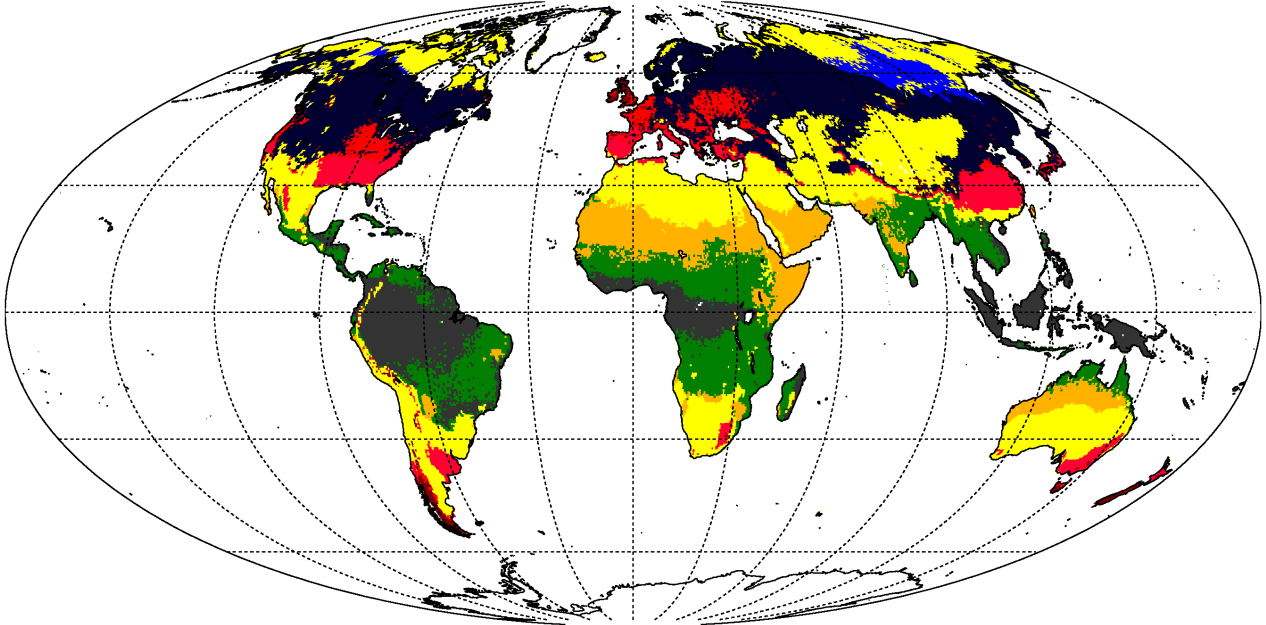
Species composition



- Polar desert
- Arctic/alpine tundra
- Desert
- Arid shrubland/steppe
- Xeric woodlands/scrub
- Short grassland
- Tall grassland
- Dry savannas
- Moist savannas
- Tropical deciduous Forest
- Tropical rain forest
- Tropical seasonal forest
- Temperate broad-leaved evergreen forest
- Temperate deciduous forest
- Temperate conifer forest
- Temperate/boreal mixed forest
- Boreal evergreen forest/woodland

Data-based map of potential natural vegetation (Haxeltine and Prentice, 1996)

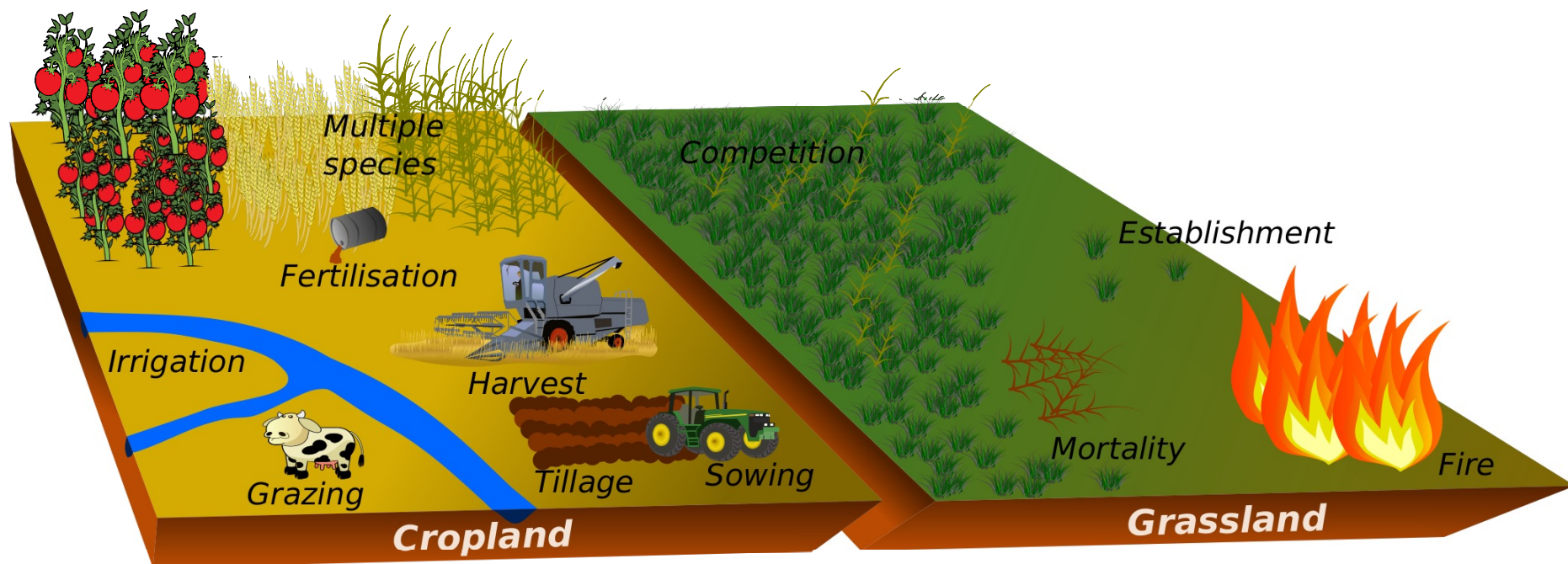
LPJ-GUESS dominant potential vegetation type 1981-2000



- Boreal needle evergreen
- Boreal needle evergreen (SI)
- Boreal needle summergreen
- Temp. broad summergreen
- Temp. broad summergreen (SI)
- Temp. broad evergreen
- Trop. broad evergreen
- Trop. broad evergreen (SI)
- Trop. broad raingreen
- C3 Herbaceous
- C4 Herbaceous
- Barren

Managed land: Agriculture

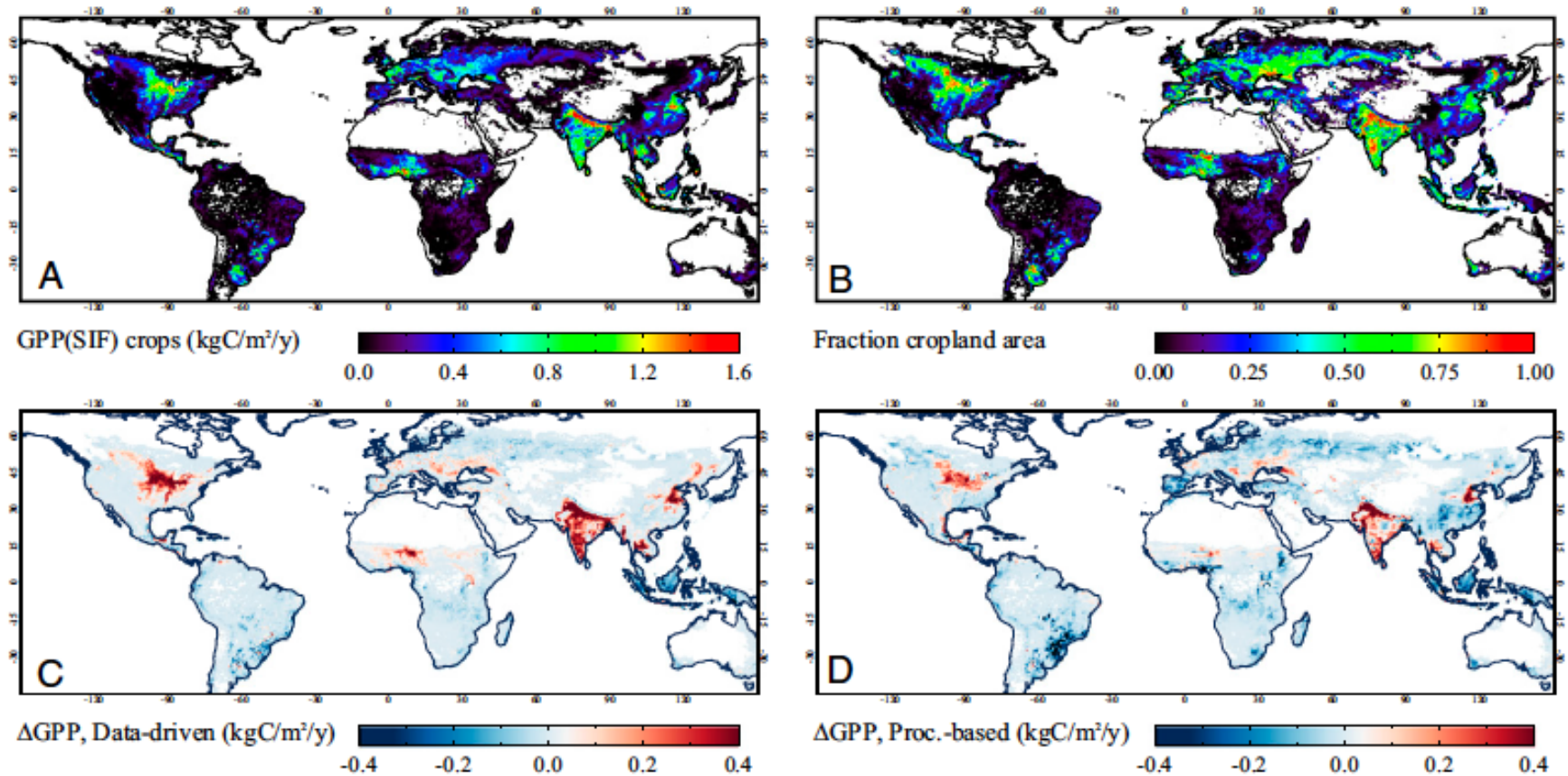
Until recently, ESMs and DGVMs concentrated on PNV.
But managed systems differ fundamentally in many respects:



Models are now being expanded to account for the range of processes existing in managed systems, e.g. agriculture.

Managed land: Agriculture

Productivity of croplands can be very different

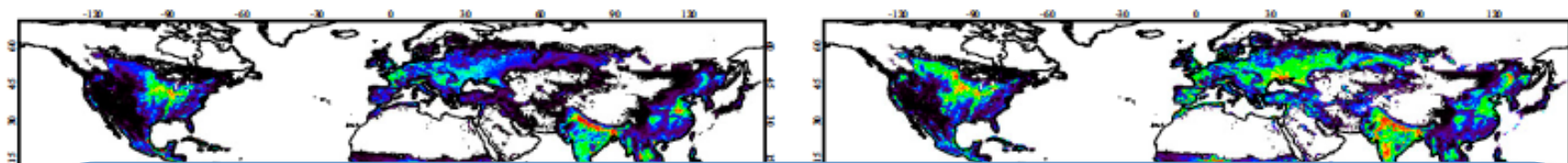


Guanter et al. (2014)

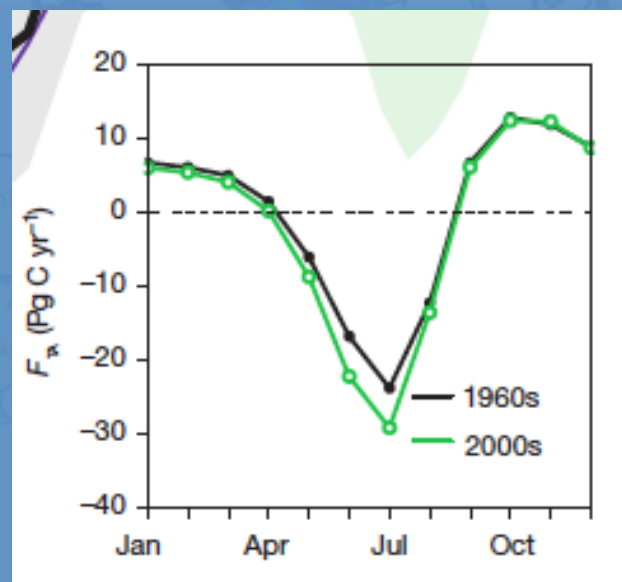
Extraction of global croplands from chlorophyll fluorescence data shows that GPP can be much higher than indicated by models which do not account for management (Guanter et al., 2014)

Managed land: Agriculture

Productivity of croplands can be very different



Recent work has shown that croplands substantially modify the seasonal cycle of atmospheric $[CO_2]$ (Zeng et al., 2014; Gray et al., 2014)



Zeng et al. (2014)

-0.4 -0.2 0.0 0.2 0.4

-0.4 -0.2 0.0 0.2 0.4

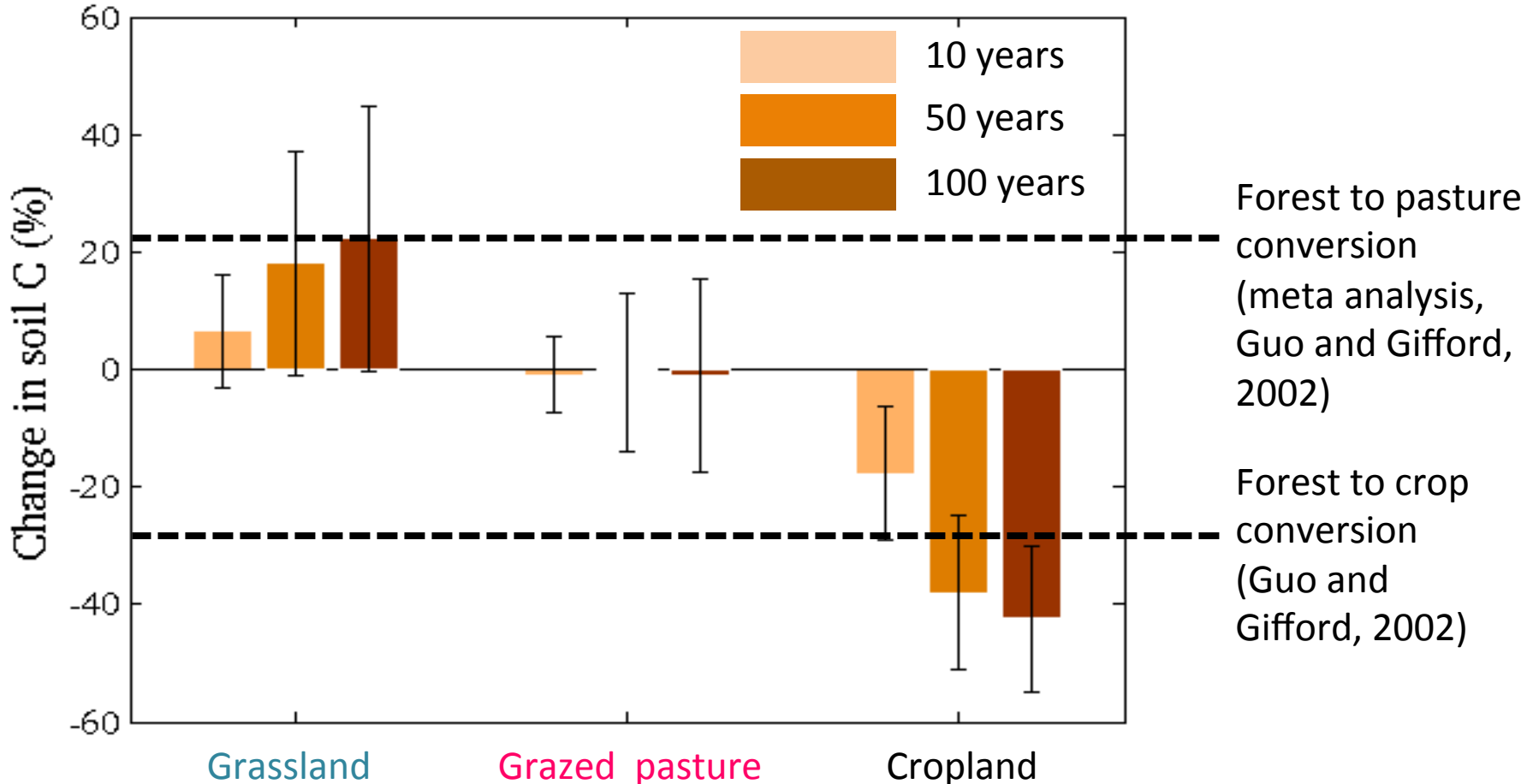
Guanter et al. (2014)

Extraction of global croplands from chlorophyll fluorescence data shows that GPP can be much higher than indicated by models which do not account for management (Guanter et al., 2014)

Managed land: Agriculture

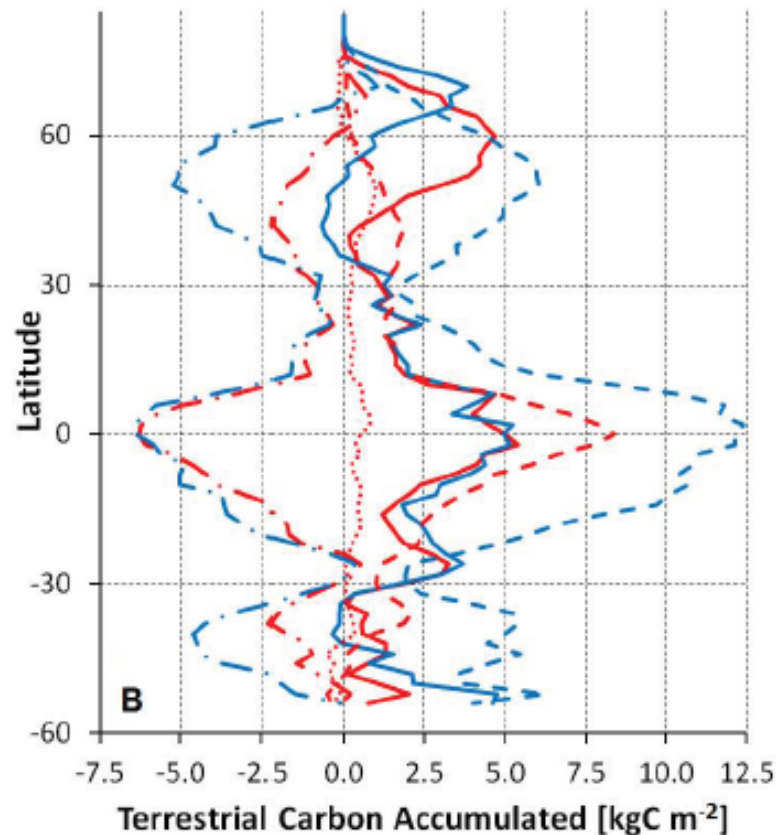
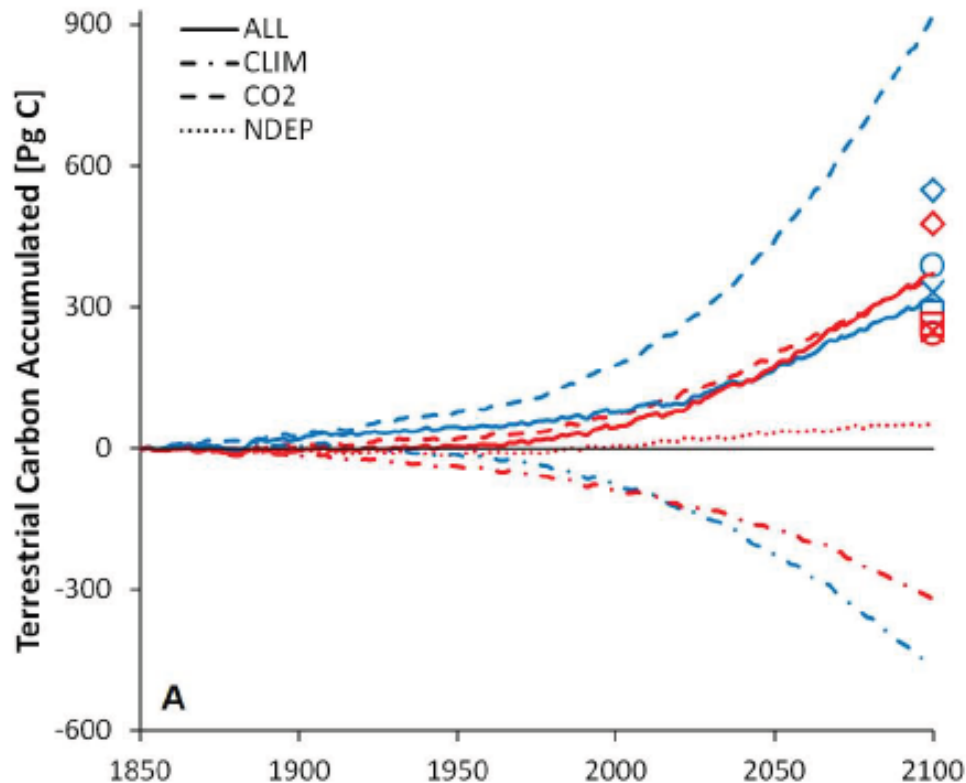
However, these productivity increases may not propagate to increases in terrestrial carbon stocks because of processes such as harvest and tillage

Change in soil carbon stocks after complete conversion from natural vegetation (global mean values):



Projecting the global carbon cycle

Change in terrestrial C accumulation:
(LPJ-GUESS, forced by MPI-ESM-LR following RCP 8.5)



— C-only version

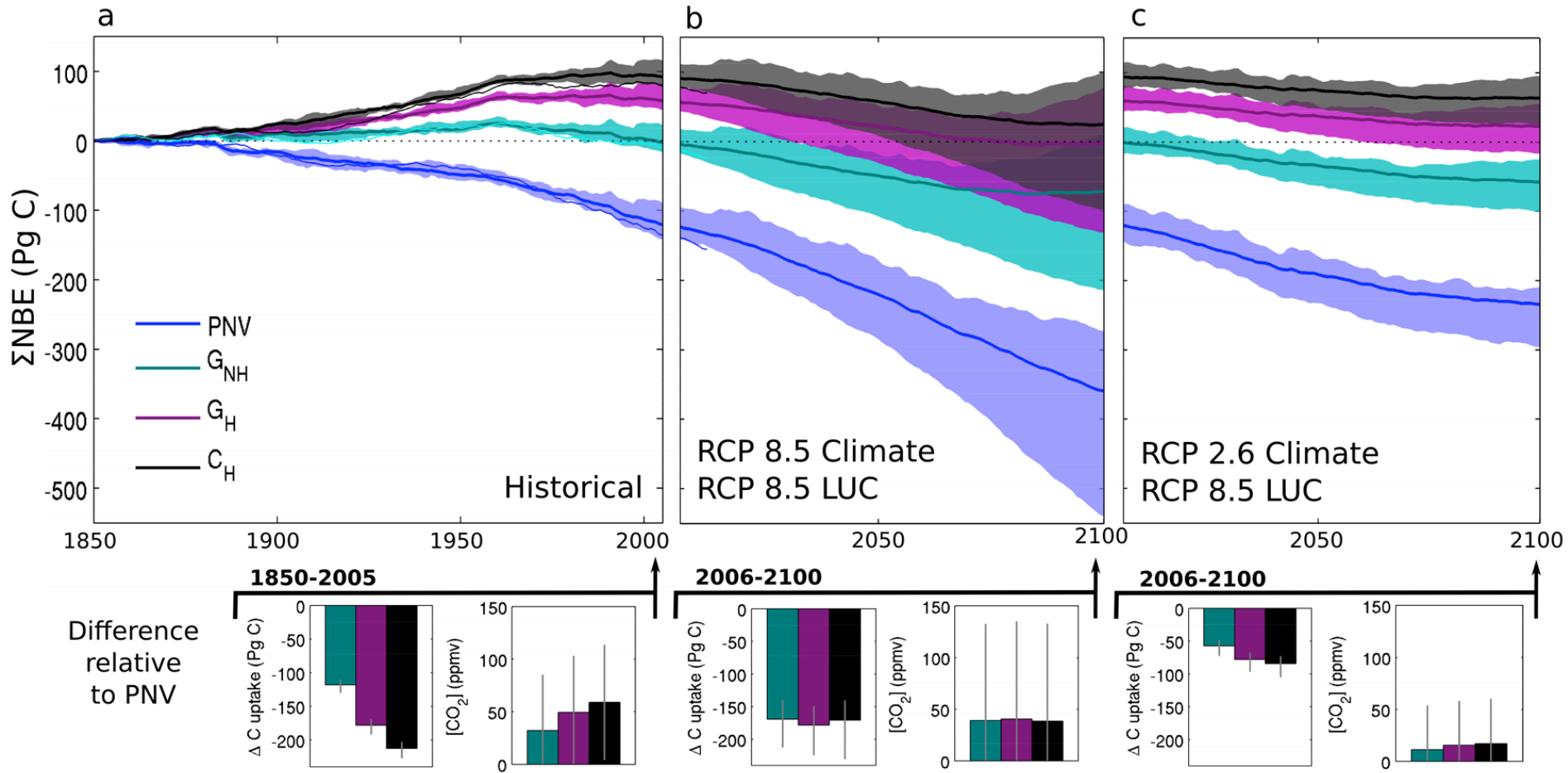
— C-N version

Symbols show other models (note opposite response)

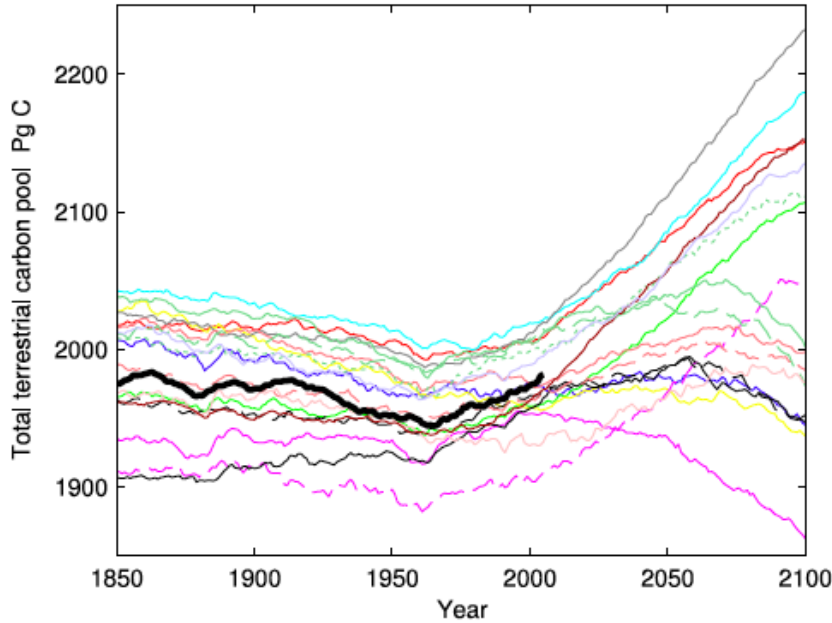
Wårlind et al. (2014)

Projecting the global carbon cycle

Together agricultural processes can make a huge difference to projections of global carbon uptake



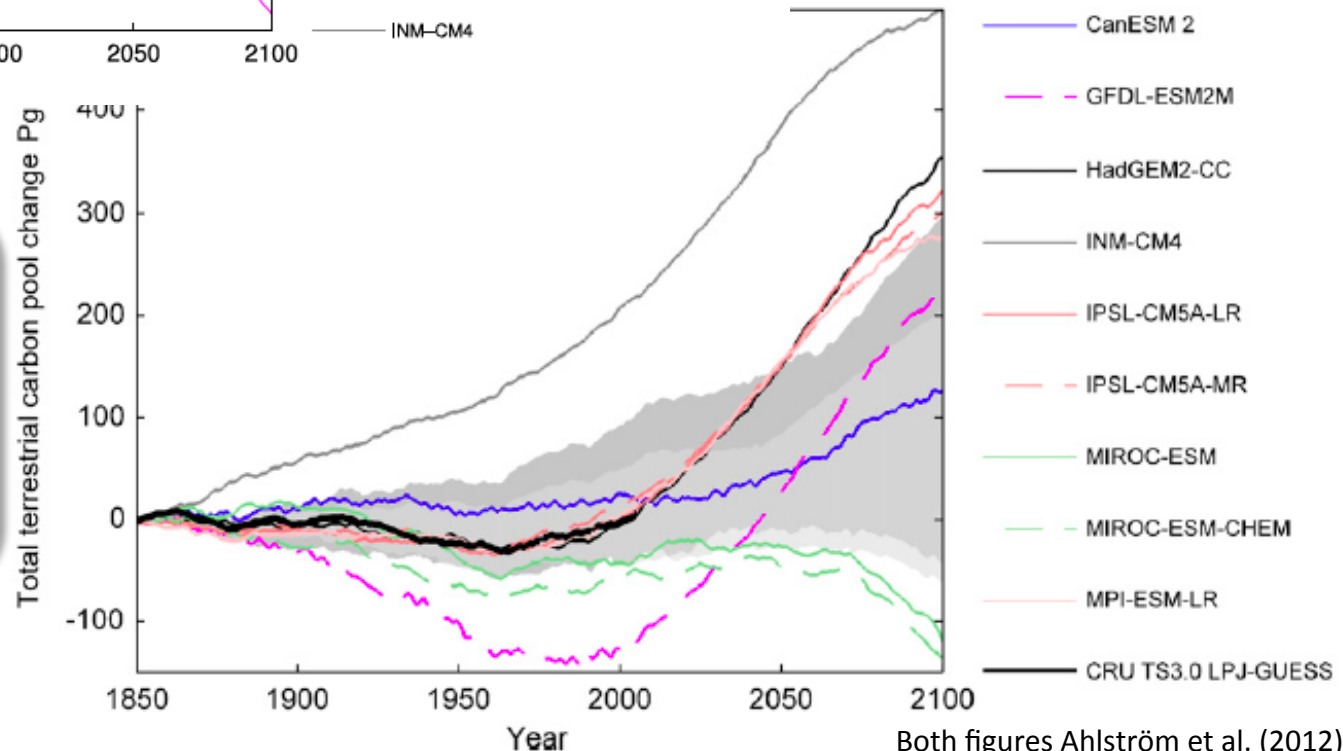
Projecting the global carbon cycle



- CanESM 2
- CCSM4
- CNRM-CM5
- FGOALS-s2
- GFDL-CM3
- GFDL-ESM2M
- GISS-E2-R
- HadGEM2-CC
- HadGEM2-ES
- INM-CM4
- IPSL-CM5A-LR
- IPSL-CM5A-MR
- MIROC-ESM
- MIROC-ESM-CHEM
- MIROC5
- MPI-ESM-LR
- MRI-CGCM3
- NorESM1-M
- CRU TS3.0

Total terrestrial carbon uptake, as calculated by LPJ-GUESS forced by a range of GCM climates (RCP 8.5)

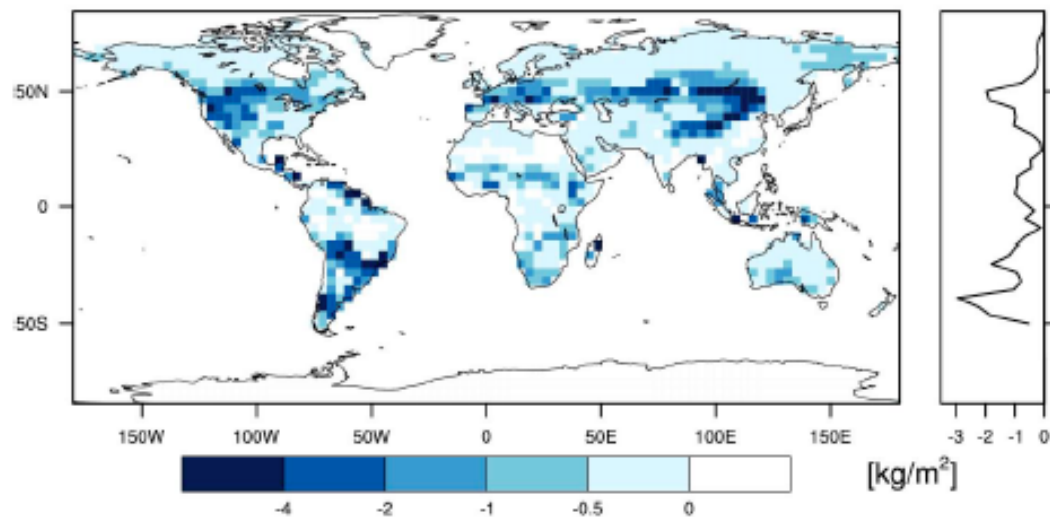
Total terrestrial carbon uptake, as calculated by several different CMIP5 ESMs (RCP 8.5)



Both figures Ahlström et al. (2012)

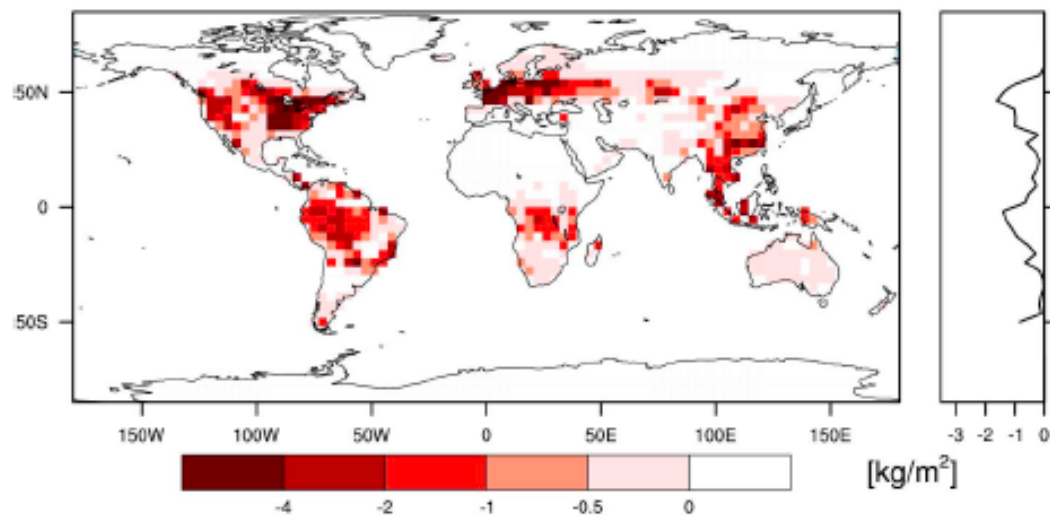
Projecting the global carbon cycle: Nutrient limitation

Nutrients other than nitrogen may also be important.



Reduction in 2070-2099 carbon storage due to nitrogen limitation in JSBACH (Goll et al., 2012)

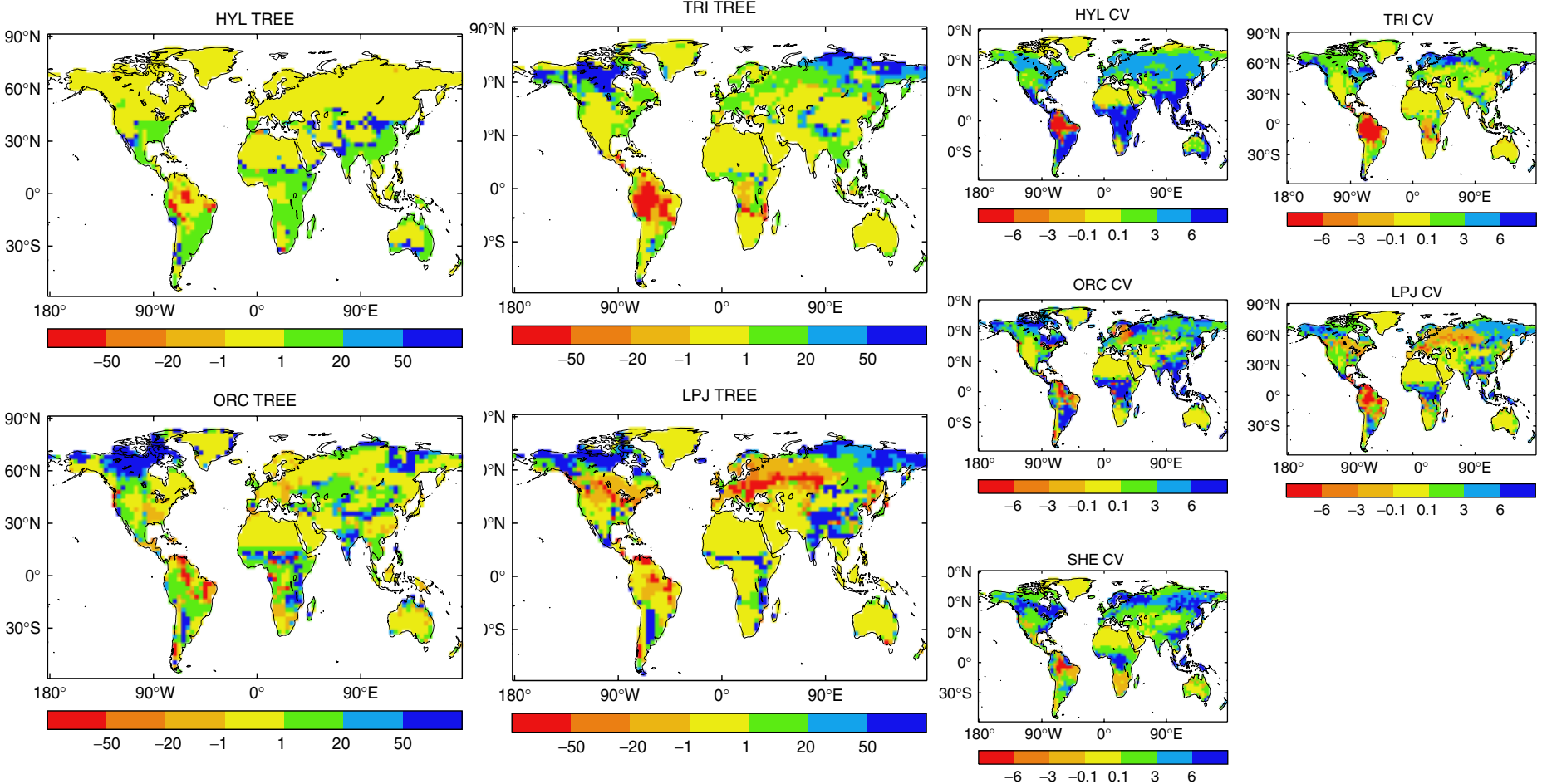
And due to phosphorus limitation (Goll et al., 2012)



Phosphorus may be the main limiting nutrient in tropical forests. Almost totally neglected in model projections, yet tropics drive expected C uptake

Projecting the global carbon cycle

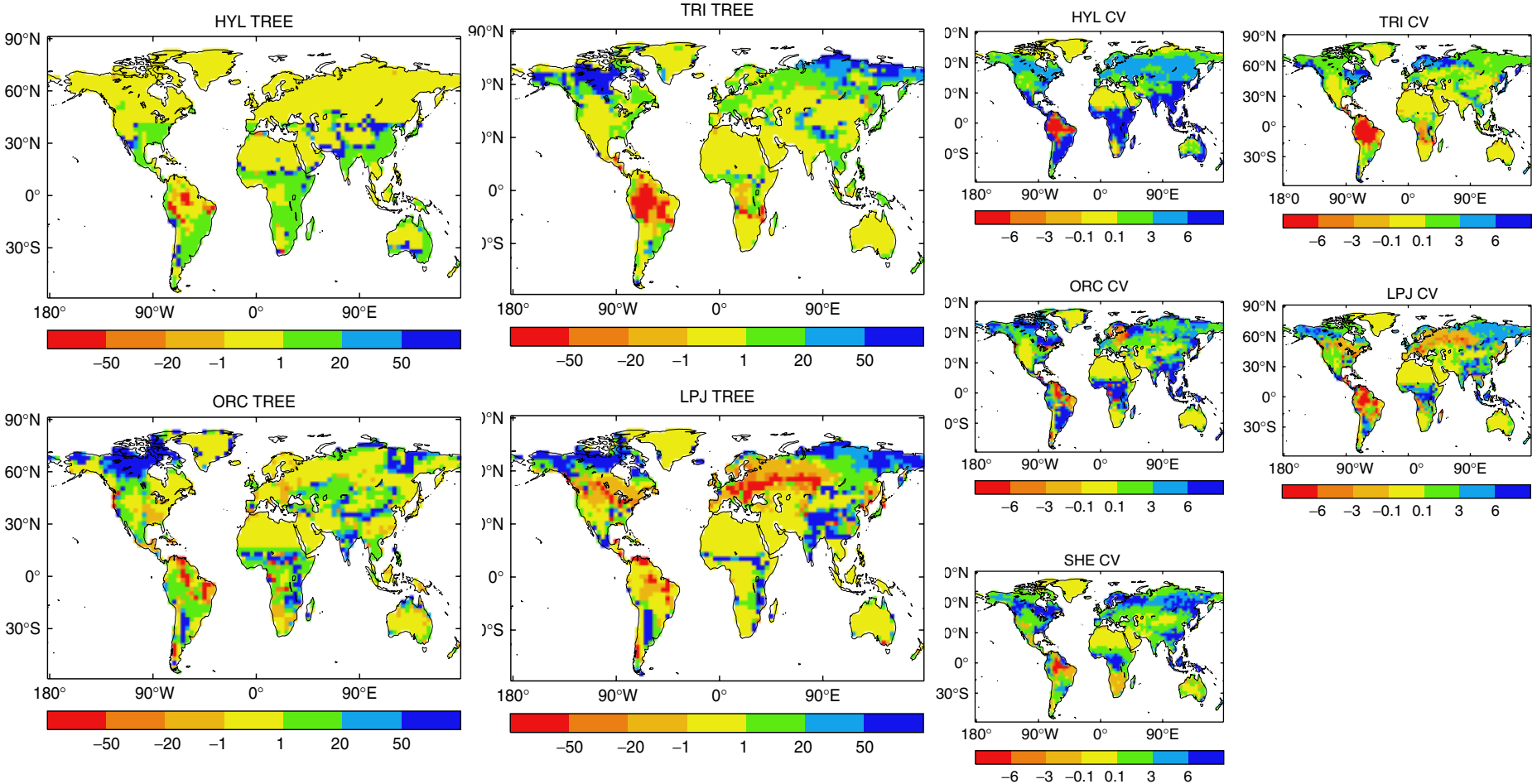
Spatial projections very different between models



Difference 1860 to 2099. Driven with climate-carbon-cycle coupling.

Projecting the global carbon cycle: Mortality

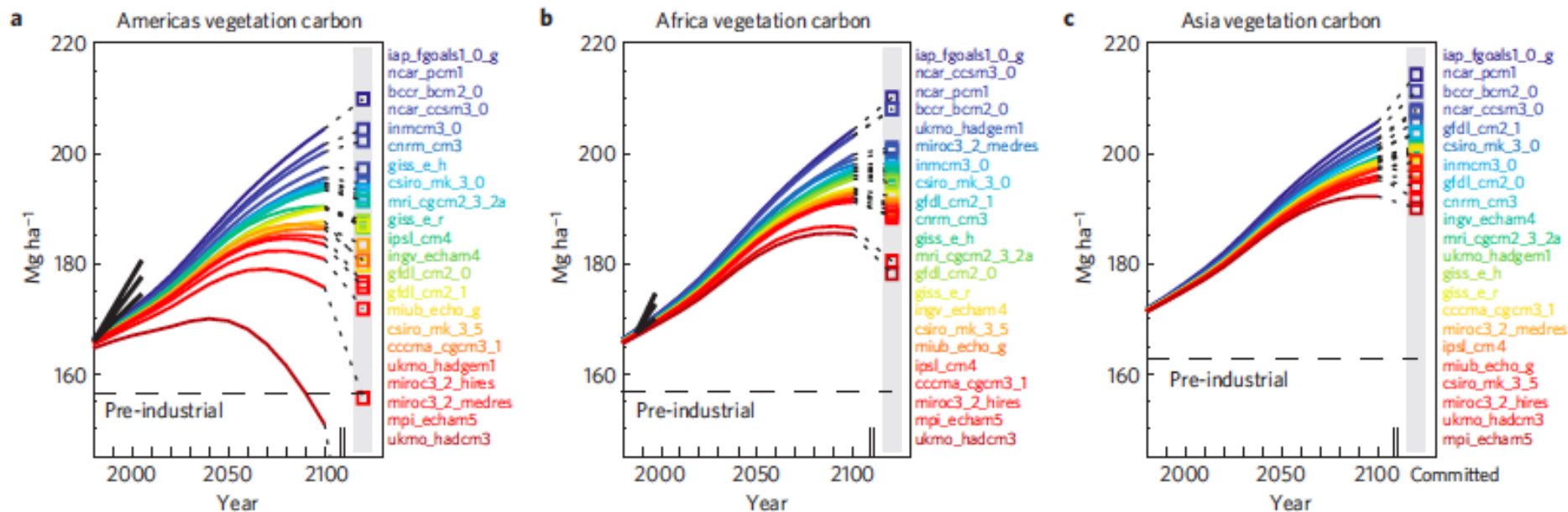
Amazon dieback?



Difference 1860 to 2099. Driven with climate-carbon-cycle coupling.

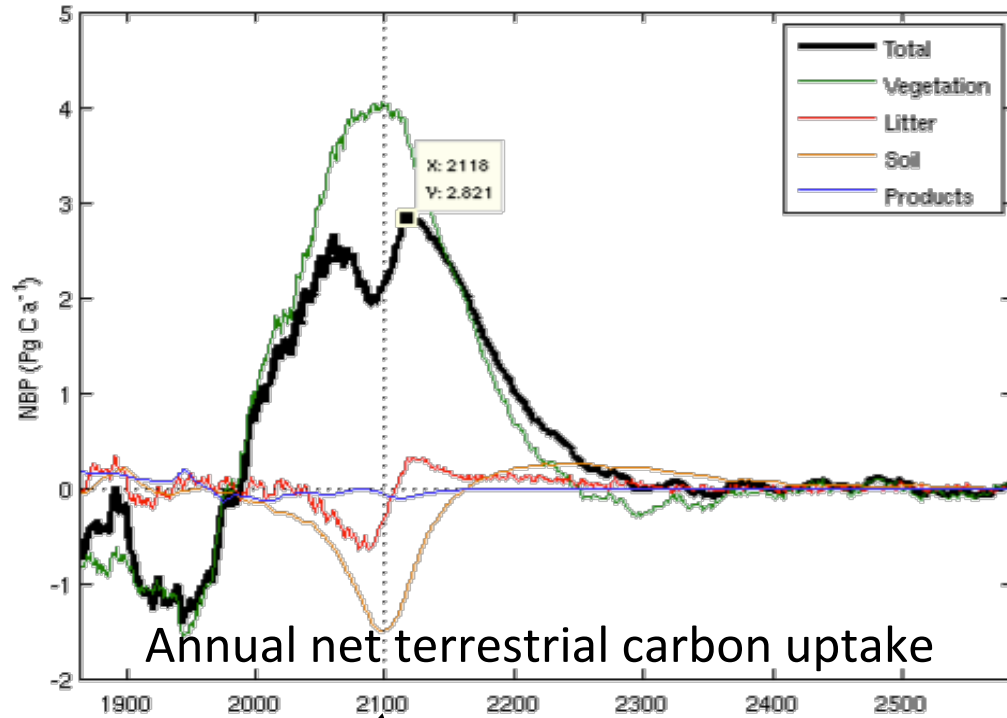
Projecting the global carbon cycle: Mortality

Or robust tropical forests?



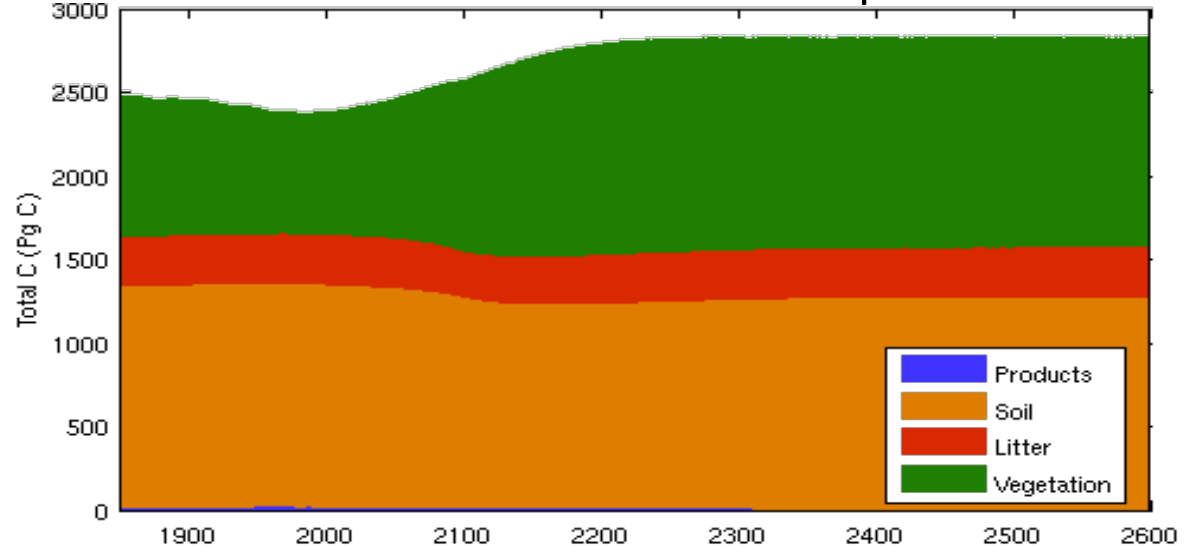
TRIFFID model forced with climates from a range of GCMs. Only one climate realisation resulted in a loss of carbon.

Projecting the global carbon cycle: Lags



Changes in many aspects of the terrestrial biosphere are not quasi-instantaneous. They show significant lags. Whilst photosynthesis and respiration rates may respond rapidly, vegetation dynamics and soil carbon pools adjust over much longer timescales.

Total terrestrial biospheric carbon



Summary and implications

- A lack of appropriate observations means that we must rely on models to understand the overall role of the land biosphere in the carbon cycle.
- These models attempt to capture the key processes, but in order to be applied at the global scale they adopt highly simplified process representations
- Models generally project a strong take up of carbon by the terrestrial biosphere under environmental change. However, they don't universally agree on this, and there are many uncertainties, relating to both explicitly resolved processes and to missing processes (e.g. peatlands).
- As you will see in the coming talks, these changes in natural vegetation properties under environmental change mean that calculations of human impacts based on the current land system may not hold in the future.
 - E.g. The tropical forest may be a much larger C store in the future, making it even more important than it is now for keeping C out of the atmosphere. Or it may dieback anyway, meaning that today's deforestation emissions would occur naturally in the future anyway.

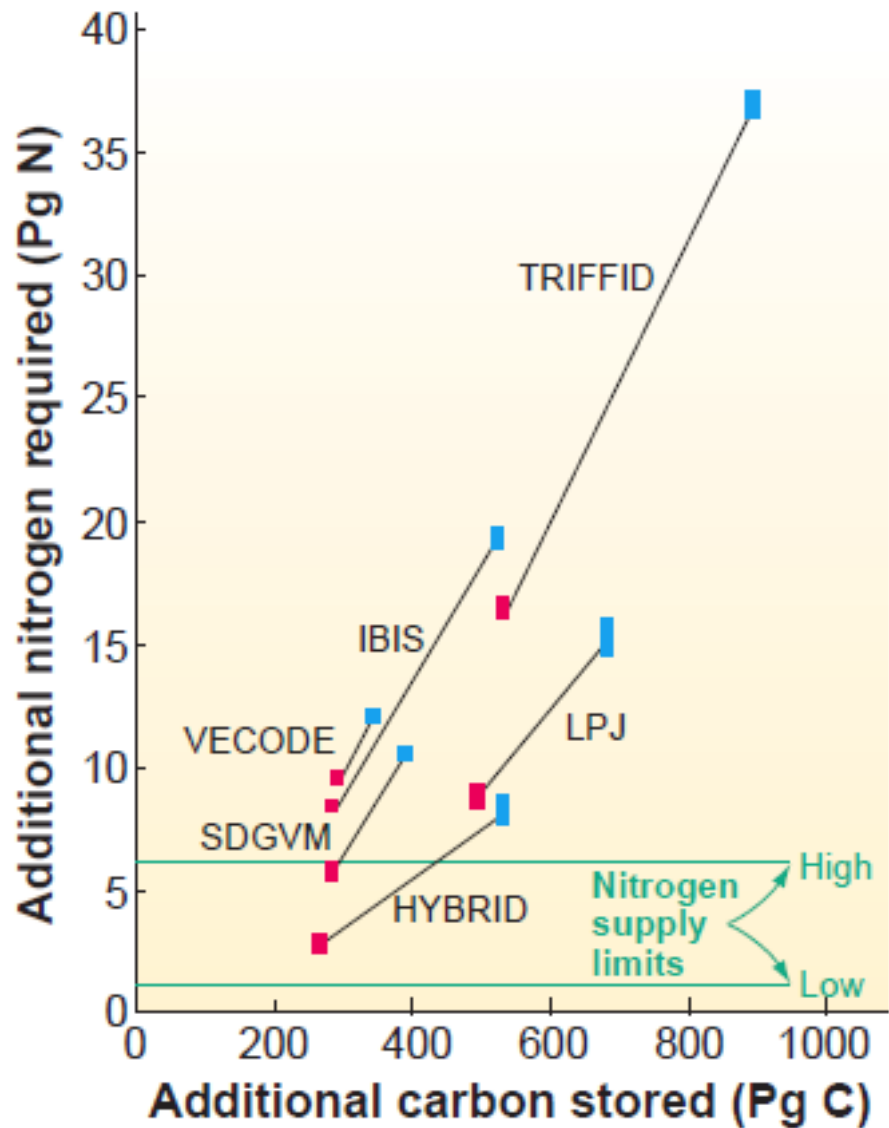
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Projecting the global carbon cycle: Nutrient limitation

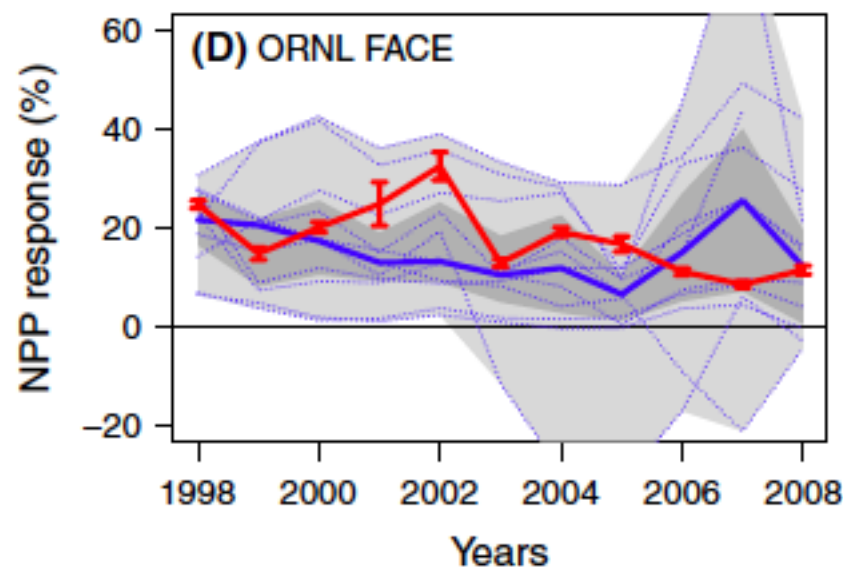
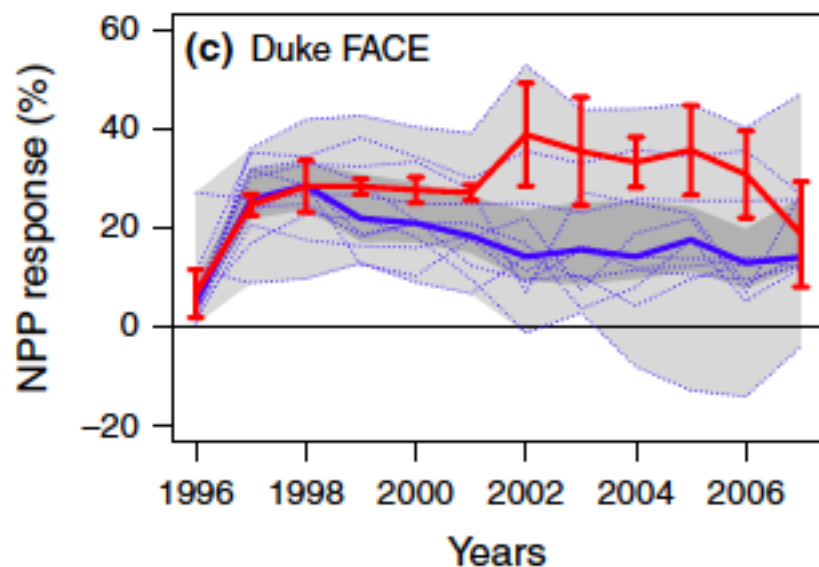


Must take results from models without nutrient constraints with a pinch of salt...

Projecting the global carbon cycle: Nutrient limitation

FACE experiments differ their long-term response to elevated $[CO_2]$

Must take results from models without nutrient constraints with a pinch of salt...



— observed ■ interquartile model range ··· individual models
— multi-model mean ■ model range

Projecting the global carbon cycle: Feedbacks

Biogeochemical feedbacks

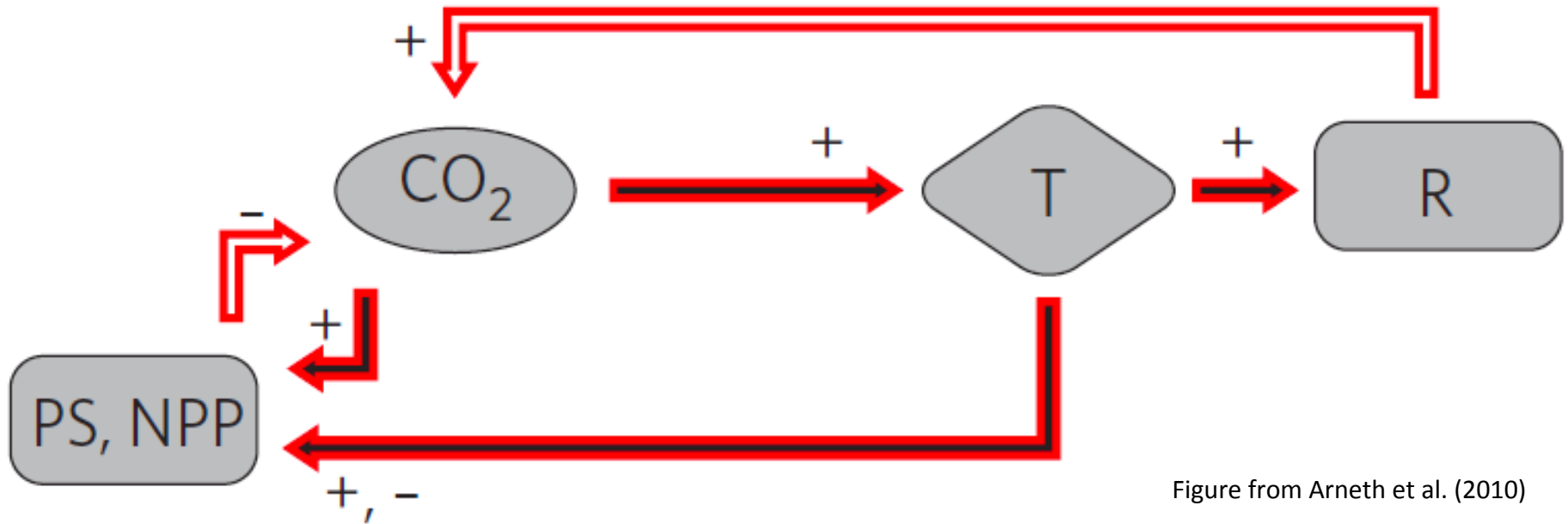


Figure from Arneth et al. (2010)

Many other feedbacks exist
Further reading: Arneth et al. (2010)