

CDOM retrieval using measurements of downwelling irradiance

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ABSTRACT

As it can strongly influence the availability of light and thus primary production, coloured dissolved organic matter (CDOM) affects the function of lake ecosystems. Therefore reliable methods are required for the monitoring of CDOM concentration. A new method using downwelling irradiance was tested for applicability in four selected lakes of the Bavarian Osterseen Lake District, which consists of 19 naturally connected freshwater lakes of different trophic level. The method separates between the direct and diffuse part of the incident light in order to handle the strong variability of the underwater light field. It is implemented in the software WASI, which is capable to retrieve water constituents by inverse modeling. During field campaigns downwelling irradiance measurements using RAMSES sensors were made in different depths. Simultaneously, water samples were taken in three depths (0.5 m, 2 m and Secchi disk depth), from which the absorption coefficient of CDOM, a_y , was derived in the range from 190 to 900 nm using photometric absorption measurements. Concentration (defined as a_y at 440 nm) ranged from 0.33 to 1.55 m^{-1} with a mean of 0.71 $\text{m}^{-1} \pm 0.04 \text{ m}^{-1}$, the spectral slope at 440 nm from 0.0120 to 0.0184 nm^{-1} with a mean of 0.0145 $\pm 0.0008 \text{ nm}^{-1}$. These laboratory measurements from water samples were compared to CDOM concentration obtained by inverse modeling of downwelling irradiance measurements using WASI. For sensor depths lower than 1 to 1.5 m large uncertainties were observed. The measurements in 2 m depth and at Secchi disk depth yielded good correlation between water sample and WASI derived data ($R^2 = 0.87$) with a mean standard deviation of 0.06 m^{-1} for the determined CDOM concentrations. This new method is an alternative to laboratory analysis of water samples from *in situ* measurements of CDOM concentration.

Keywords: inland waters, monitoring, water constituents, CDOM, downwelling irradiance, RAMSES sensors, inverse modeling, WASI

1. INTRODUCTION

Coloured dissolved organic matter (CDOM) influences the structure and function of aquatic ecosystems¹ like freshwater lakes. Comprising the chromophoric fraction of the dissolved organic carbon (DOC), CDOM is a strong absorber of the incident radiation especially in the UV and the blue spectral region. Therefore optical properties of CDOM contribute to the inherent optical properties (IOPs) of lake ecosystems² and significantly affect the availability of light and thus transparency³. At high concentration CDOM negatively influences underwater photosynthetic active radiation (PAR, 400 – 700 nm) and light can get the limiting factor for photoautotrophic algae⁴⁻⁶. On the other hand, CDOM is required as an important carbon source for heterotrophic bacteria in freshwater lakes⁷ and plays a central role in the mobility of nutrients and different organic and inorganic toxins. As major parts of CDOM enter lake ecosystems via input of riverine water and surface water runoff, monitoring of CDOM concentration and type can give important information about changes in the surrounding catchment area⁸. Like all aquatic ecosystems, freshwater lakes are exposed to numerous environmental changes mostly of anthropogenic nature, changes that could also alter natural carbon fluxes and consequently structure and function of freshwater lakes¹. For instance, increasing CDOM concentration was observed in rivers draining peatlands in the northern hemisphere during the last decades^{9, 10}. The driving factor is still under discussion¹¹. The hypotheses include rising temperature as consequence of climatic change, changing river discharge, and induced changes in the landscape water balance. Methods for the retrieval of CDOM are therefore of great interest to predict future changes in carbon cycle and possible changes in water quality¹¹.

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Hyperspectral *in situ* radiation measurements could be a proficient method for spatial CDOM monitoring. Determination of CDOM has been successfully performed by analysis of spaceborne¹¹ and airborne¹² images, but these methods are bound to several limitations. Multispectral spaceborne sensors have only broad spectral bands and data acquisition is restricted to the fixed times of the satellite overpass, whereas hyperspectral airborne sensors offer the required spectral resolution but are too cost intensive for regular observations. In addition, due to clouds and sunglint many images cannot be used. With *in situ* radiation measurements samples can be taken any time and throughout the water column. The use of reliable bio-optical models for the inversion of radiation measurements offers a much more efficient data analysis than laboratory. Additionally *in situ* radiation measurement can contribute to interpretation and validation of remote sensing algorithms¹³.

A new method using under water measurements of downwelling irradiance, E_d , turned out to be an effective method for the determination of sensor depth¹³ and phytoplankton concentration¹⁴. Treating the direct and diffuse components of E_d separately, this new model can correct effects of disturbance from the water surface. Being of great interest for the monitoring of future changes in carbon cycle, the new model is tested for the retrieval of CDOM concentration.

2. IRRADIANCE MODEL

The downwelling irradiance in water is rarely used to derive water constituents since the measurements are generally highly variable (Figure 1a). The variability is mainly caused by the water surface¹⁵. Waves, ripples and foam alter steadily the refraction angle of the incident rays, leading to focusing and defocusing effects^{16, 17}. The induced changes of intensity are typically in the order of 20 to 40 % in the upper few meters^{18, 19}, but flashes can be an order of magnitude above average²⁰. Because the incident angles are different for the direct (E_{dd} , representing the sun disk) and diffuse (E_{ds} , representing the sky) irradiance components, the effect of the water surface on the two components is different, i.e. the intensity changes of E_{dd} and E_{ds} are only weakly correlated¹⁵. Consequently, the spectral shape of E_d is fluctuating as well (red curve of Figure 1a) since the spectral shapes of E_{dd} and E_{ds} are different (Figure 1b). A model that can handle this variability of E_d intensity and spectral shape was developed recently^{21, 22}. The concept is summarized below, for details see the cited references.

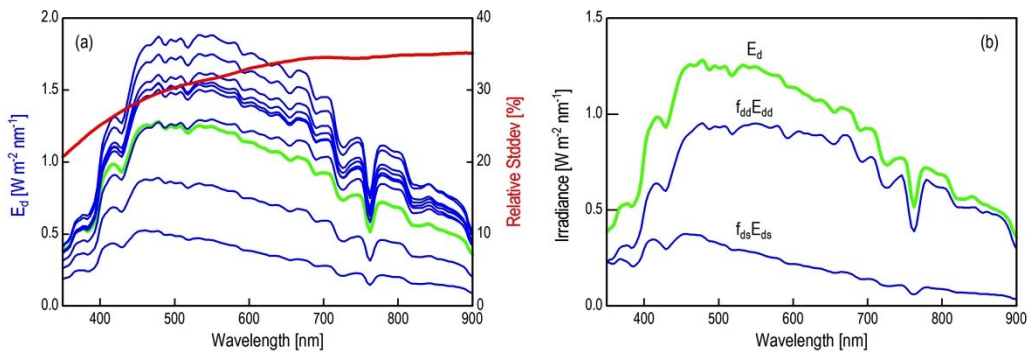


Figure 1. Illustration of downwelling irradiance in water. (a) Short-term variability of E_d . 10 measurements were made within 100 s just below the water surface. The standard deviation illustrates the associated changes of spectral shape. (b) Differences between the irradiance components. The described model was used to separate the direct (E_{dd}) and diffuse (E_{ds}) components for the green marked E_d measurement of panel (a). The relative intensities were $f_{dd}=0.97$ and $f_{ds}=0.69$.

The downwelling irradiance at depth z and wavelength λ is expressed as a weighted sum of E_{dd} and E_{ds} as follows:

$$E_d(\lambda, z) = f_{dd}E_{dd}(\lambda, z) + f_{ds}E_{ds}(\lambda, z) \quad (1)$$

The wavelength-independent parameters f_{dd} and f_{ds} describe the intensities of E_{dd} and E_{ds} relative to conditions with a plane water surface. They change from measurement to measurement and are not known *a priori*. f factors <1 correspond to intensity decrease, f factors >1 to intensity increase. They can be determined during data analysis if (1) a model for calculating $E_{dd}(\lambda, z)$ and $E_{ds}(\lambda, z)$ for a plane water surface exists, and (2) the dimensionality of the E_d measurements is higher than the number of model parameters, i.e. wavelength range, number of channels and signal-to-noise-ratio are sufficiently high.

A depth-dependent model of $E_{dd}(\lambda, z)$ and $E_{ds}(\lambda, z)$ has been developed by Gege²¹. It approximates the downwelling irradiance components at the water surface (at $z = 0$) using the model of Gregg and Carder²³, which was extended concerning spectral range and spectral resolution. Parameters are the Sun zenith angle (θ_{sun}) and the atmospheric properties Angström exponent of aerosols (α), turbidity coefficient (β), scale height of ozone (H_{oz}) and scale height of precipitable water (WV). α is a measure of the aerosol type; β , H_{oz} and WV are measures of concentration. The four atmospheric parameters α , β , H_{oz} and WV are usually unknown, but can be estimated by inverse modeling of an above-water measurement of $E_d(\lambda)$.

The downwelling irradiance components at depth $z > 0$ are calculated using the Lambert-Beer-Law $E(z) = E(0) \exp[-Kz]$. In contrast to the widely used approach of applying the law to E_d , it is applied here to E_{dd} and E_{ds} separately in order to account for the different path lengths of the rays forming $E_{dd}(z)$ and $E_{ds}(z)$. This separation reduces significantly the model complexity for the attenuation coefficient K and improves accuracy of $E_d(z)$ calculation with increasing depth. While K_d of the usual approach is an apparent optical property (AOP) with complicated dependency on geometry parameters²⁴, K_{dd} and K_{ds} of the new approach can be expressed as products of $u(\lambda) = b_b(\lambda) / [a(\lambda) + b_b(\lambda)]$, which is an inherent optical property (IOP), and path length normalized to z . The normalized path length is very close to 1 for E_{dd} , and proportional to the cosine of the underwater Sun zenith angle for E_{ds} ²¹. The absorption coefficient $a(\lambda)$ and the backscattering coefficient $b_b(\lambda)$ are the optical properties of the water and comprise all extractable information about the water constituents. The irradiance model can exploit the information of $a(\lambda)$, but not reliably that of $b_b(\lambda)$ ¹⁴. This enables the determination of CDOM, which is a strongly absorbing component.

3. METHODS

3.1 Study site

The Osterseen Lake District (47°48'N, 11°18'E) developed after the last glacial epoch in the Alpine foothills in the South of Munich (Bavaria, Germany). Consisting of 19 naturally connected freshwater lakes (Figure 2), the district can be divided into freshwater lakes located in the North, South and East of the district. Resulting from a contamination of the lake supplying groundwater layer due to municipal wastewater and agricultural management of the catchment area the Osterseen Lake District, originally oligotrophic, was exposed to eutrophication. Because of their connection a nutrient gradient from south to north, following the natural incline of the territory, developed over the years. Being of similar age, origin and morphology contemporary with exposure to the same climatic environment, the freshwater lakes only differ in groundwater supply, trophic level and connectivity to surrounding dystrophic bogs as potential CDOM source. Therefore the Osterseen Lake District contains a wide spectrum of different lake types, optimal for comparative limnological studies²⁵.

3.2 Field campaign data

Water samples were taken from June to September 2012 in a two-weekly pattern at five sampling sites (Figure 2) selected in the Lake Ostersee (OS1, OS2), Lake Eishaussee (ES1), Lake Fohnsee (FS1) and Lake Sengsee (SS1). Regular sampling was performed only in the southern and eastern lakes of the district, since the northern lakes can't be reached by boat directly. To cover also the northern lakes, 14 of them were sampled during a single campaign on August 14, 2012. Due to its size (1,18 km²) two sampling sites were chosen in Lake Ostersee. At each test site and each campaign, three replicates of one litre water were sampled in three depths (0.5 m, 2 m and Secchi disk depth (SDD)) using a standard water sampler according to Ruttner (Hydro-Bios). If Secchi disk depth was 2 m or less, the second sample was taken at 1 m instead of 2 m water depth. In the laboratory, water samples were filtered using 0.45 µm pore-size cellulose-acetate filters (Sartorius Stedim Biotech), filled in PE-bottles and stored in a fridge in dark. Further analysis was performed within one month after the campaign. Then samples were analysed for absorption using a CARY double-beam UV-VIS spectrometer (Agilent Technologies) following the method of Gege²⁶. The transmission of each sample was measured twice in quartz cuvettes with 5 and 10 cm pathlength (Starna Spectrosil Cuvettes 1/Q/50, 1/Q/100) from 190 to 900 nm at 1 nm intervals. The absorption spectrum of the sample was determined from the ratio of the two transmission spectra using the Lambert-Beer law. The CDOM absorption spectrum, $a_{\lambda}(\lambda)$, was calculated by subtracting the same pure water absorption spectrum that is also used in the model. No correction of offset or temperature was performed. Concentration of CDOM is defined as absorption coefficient at 440 nm.

Simultaneously to water sampling, the downwelling irradiance was measured in water using a hyperspectral TRIOS RAMSES ACC-VIS irradiance sensor (wavelength range 320-900 nm, spectral sampling interval 3.3 nm), calibrated by

TRIOS in 2008. A preferable clear sky and a sun elevation of at least 30° ²⁷ were chosen for the measurements. Mounted on a metal frame the upward facing sensor was lowered into the water column using a cable winch. As recommended in literature¹³ data collection started with a measurement above water surface $E_d(0+)$, followed by a measurement just below water surface $E_d(0-)$. Then the sensor was lowered beneath Secchi disk depth and moved upwards in 0.5 m steps. Measurements were made up to depths of 4.5 meters. The water depth was recorded simultaneously to the E_d measurements by a pressure sensor attached to the RAMSES device. For calibration and equidistant resampling in a wavelength grid of 320 to 950 nm with a spectral resolution of 3.3 nm the software MSDA_XE 8.5 (TriOS GmbH) was used.

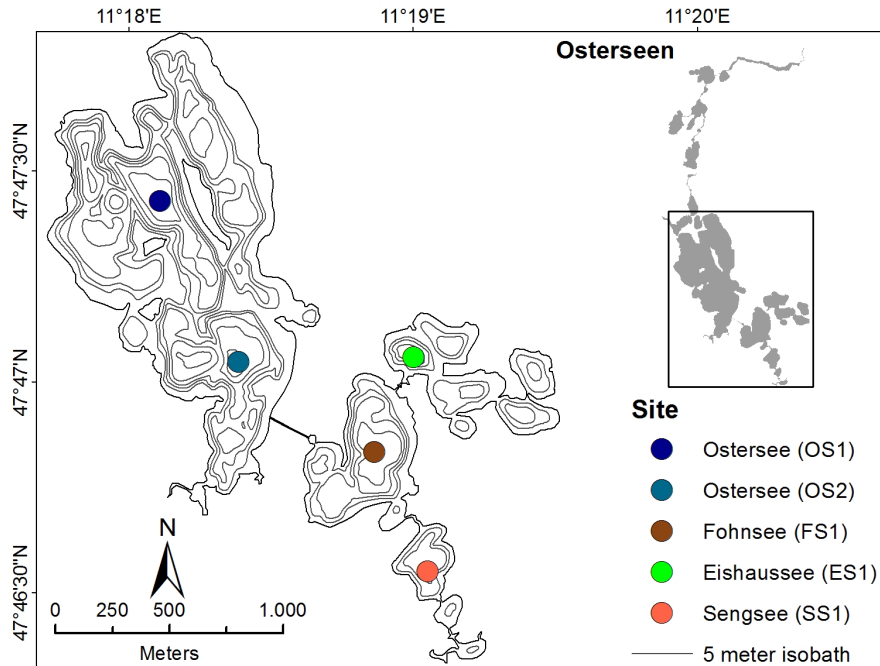


Figure 2. Map of the 19 connected freshwater lakes of the Osterseen Lake District (right) situated in the South of Munich (Bavaria, Germany). Points display the five sampling sites (OS1, OS2, ES1, FS1 and SS1) each at maximum depth in the four selected lakes of the district (left), Lake Ostersee (30.6 and 27.7 m), Lake Eishaussee (20.1 m), Lake Fohnsee (24.6 m) and Lake Sengsee (15.4 m). Gray lines show the 5 meter isobaths.

3.3 Data analysis

The Water Color Simulator WASI was used for data analysis. It allows to simulate and analyse different types of spectral measurements of instruments disposed above the water surface and submerged in the water. Data analysis is done by inverse modeling. The described irradiance model is implemented in WASI version 4. For details of its parameterisation and inverse modeling algorithm see Gege²¹.

The focus of this study is CDOM retrieval. As described in section 2, CDOM can be determined from irradiance measurements by exploiting the information that is implicitly included in the absorption coefficient of the water body, $a(\lambda)$. WASI treats $a(\lambda)$ as sum of four absorption coefficients that represent pure water, detritus (non-algal particles), phytoplankton, and CDOM. For our dataset, detritus was neglected and phytoplankton was described by three specific absorption coefficients representing the dominant classes of our test sites: diatoms, dinoflagellates, and green algae. CDOM absorption is calculated in WASI as $a_Y(\lambda) = Y a_Y^*(\lambda)$, where $a_Y^*(\lambda)$ is the specific absorption coefficient, normalized at a reference wavelength λ_0 , and $Y = a_Y(\lambda_0)$ is the absorption coefficient at λ_0 in units of m^{-1} . $a_Y^*(\lambda)$ is a measure of the CDOM type, and Y of CDOM concentration. $a_Y^*(\lambda)$ is frequently approximated by an exponential function, $a_Y^*(\lambda) = \exp[-S(\lambda - \lambda_0)]$. The spectral slope S at $\lambda_0 = 440 \text{ nm}$ is typically in the range between 0.01 nm^{-1} for humic acid dominated waters and 0.02 nm^{-1} when fulvic acids prevail. WASI can use either the exponential approximation, or a spectrum $a_Y^*(\lambda)$ imported from file.

In this study, the model parameters C_i , X , Y , S , f_{dd} , f_{ds} , α , β , H_{oz} , WV and z were treated as potential unknowns for individual E_d measurements and determined by inverse modeling. The C_i are the concentrations of phytoplankton groups (C_0 : weighted sum of the five phytoplankton spectra, C_1 : cryptophyta “low”, C_2 : cryptophyta “high”, C_3 : diatoms, C_4 : dinoflagellates, C_5 : green algae), X is suspended matter concentration, Y and S are the before-mentioned CDOM parameters, f_{dd} and f_{ds} are the relative intensities of direct and diffuse irradiance as described above, H_{oz} , α , β and WV are the before-mentioned parameters of the atmosphere, and z is sensor depth.

Inverse modeling was carried out for all E_d measurements. The defaults were set to $C_0 = 2 \text{ mg m}^{-3}$, $C_i = 0$ for $i = 1..5$, $X = 2 \text{ g m}^{-3}$, $Y = 0.5 \text{ m}^{-1}$, $f_{dd} = 1$, $f_{ds} = 1$, $\alpha = 1$, $\beta = 0.200$, $H_{oz} = 0.300 \text{ cm}$, $WV = 1.50 \text{ cm}$, $z = 2 \text{ m}$. S was set to 0.0140 nm^{-1} . To test the influence of the initial values on the fit results, different initial values of Y were used (0.002 m^{-1} , 1 m^{-1} , 5 m^{-1} , 10 m^{-1}). In a next step different S values were set (0.0040 nm^{-1} , 0.0250 nm^{-1}), S was set as fit parameter with an initial value of 0.0140 nm^{-1} , and further the average specific absorption spectrum $a_{Y,OS}^*(\lambda)$ of all measurements conducted in the Osterseen Lake District was tested instead of the exponential approximation. The influence of phytoplankton absorption on CDOM retrieval was analyzed by using either C_0 alone or C_3 , C_4 and C_5 together as fit parameters, with initial values of 1 in each case.

4. RESULTS

4.1 Optical properties of CDOM in the Osterseen Lake District

A summary of the results from the regular campaigns at the lakes Ostersee, Eishaussee, Fohnsee and Sengsee is presented in Table 1. For the individual laboratory measurements, the concentration of CDOM, $a_Y(440)$, ranges from 0.33 to 1.55 m^{-1} with a mean and standard deviation (SD) of $0.71 \pm 0.04 \text{ m}^{-1}$. The spectral slope, S , ranges from 0.0120 to 0.0184 nm^{-1} with a mean and SD of $0.0145 \pm 0.0008 \text{ nm}^{-1}$.

As can be seen in Table 1, the results for $a_Y(440)$ are very similar for the two methods and agree within a SD. The SD is typically in the order of 5 % for the laboratory measurements and 10 % for the E_d fits, except for the site SS1, where SD of both methods is increased by a factor of two. The agreement of the results for S is less convenient. The comparably high SD of S derived from E_d fits indicates that S cannot be determined accurately by inverse modeling of E_d measurements (see section 4.2).

The lakes reveal significant differences both in CDOM concentration and type. In Lake Fohnsee the highest concentrations were observed, in Lake Sengsee the lowest. Lake Ostersee and Lake Eishaussee range in between. The parameter S , which is a measure of CDOM type, shows the highest values for Lake Ostersee and the lowest for Lake Eishaussee and Lake Sengsee, while Lake Fohnsee lies in between. The two sampling sites in Lake Ostersee constantly show similar values for $a_Y(440)$ and S .

Table 1. Summary of $a_Y(440)$ and S measurements at the main sampling sites. (a) Results of the laboratory measurements. (b) Results from fitting E_d spectra. N = number of measurements (different days, different depths), N' = number of single irradiance measurements.

Sampling site	(a)				(b)							
	N	$a_Y(440) [\text{m}^{-1}]$		$S [\text{nm}^{-1}]$		N	N'	$a_Y(440) [\text{m}^{-1}]$		$S [\text{nm}^{-1}]$		
		mean	SD	mean	SD			mean	SD	mean	SD	
OS1	9	0.82	0.02	0.0156	0.0004	6	120	0.73	0.06	0.0134	0.0013	
OS2	9	0.81	0.04	0.0156	0.0001	6	105	0.92	0.11	0.0137	0.0013	
ES1	15	0.62	0.04	0.1380	0.0008	9	135	0.64	0.05	0.0146	0.0012	
FS1	12	1.08	0.05	0.0145	0.0006	6	105	1.08	0.10	0.0143	0.0020	
SS1	15	0.40	0.04	0.0137	0.0011	9	135	0.42	0.09	0.0132	0.0024	

To visualize regional and temporal variabilities, the complete laboratory dataset measured at all test sites during all campaigns is illustrated in Figure 3. The earliest campaign from June 19th 2012 shows the highest $a_Y(440)$ for all sampling sites except Lake Eishaussee. Pronounced fluctuations of concentration during the summer are observed in Lake Ostersee and Lake Fohnsee, while $a_Y(440)$ in Lake Eishaussee and Lake Sengsee is more constant during summer.

Comparing the three sampling depths (0.5 m, 2 m and Secchi disk depth) no significant differences can be seen for $a_Y(440)$. The $a_Y(440)$ measurements reveal regional differences, i.e. high CDOM concentration for the northern lakes of the district (except Lake Lustsee), and relatively constant concentrations for the southern and eastern lakes. Looking at the spectral slope, the S values are similar in all lakes of the district, ranging from 0.0130 to 0.0150 nm^{-1} , with one extreme value in Lake Ostersee. A regional trend, as was observed for $a_Y(440)$, is not visible in the S data. The average of S is 0.0145 nm^{-1} , the SD 0.0008 nm^{-1} , i.e. the relative SD is 5%.

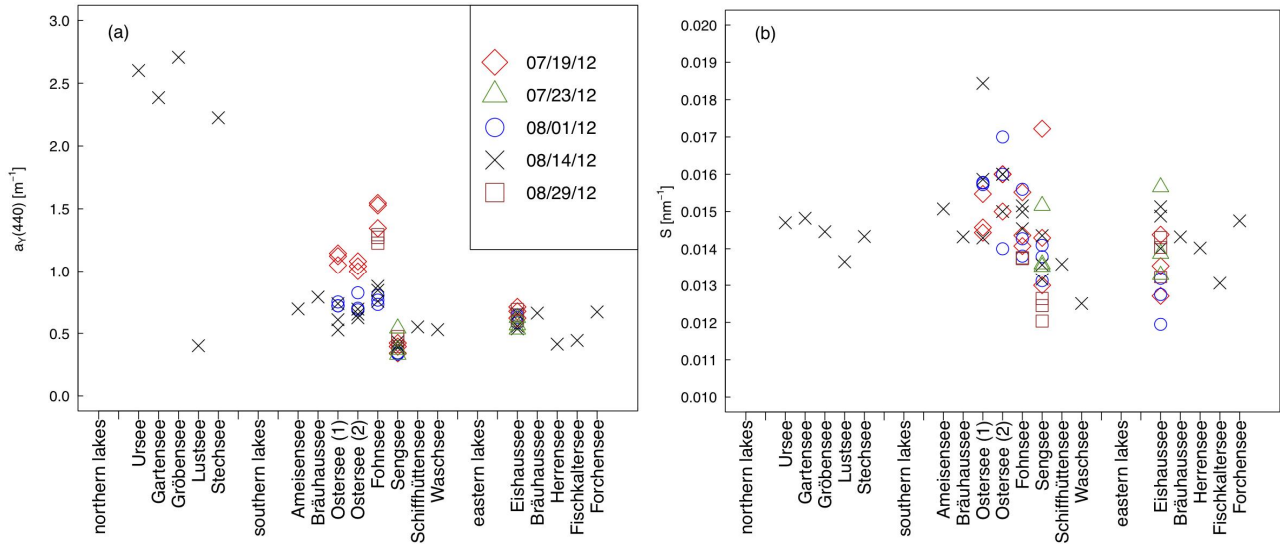


Figure 3. Regional and temporal distribution of CDOM parameters in freshwater lakes of the Osterseen Lake District. The symbols represent the five campaigns carried out during summer 2012. (a) CDOM concentration $a_Y(440)$. (b) CDOM spectral slope S .

To illustrate differences in specific absorption between the sampling sites, $a_Y^*(\lambda)$ normalized at the reference wavelength of $\lambda_0 = 440 \text{ nm}$ is presented in Figure 4. Each coefficient represents the average of all samples taken at the specified sampling site. Apparently the shape of the CDOM absorption curve is very similar for all sites. Hence, the average $a_{Y,OS}^*(\lambda)$ is taken below as input to WASI representing $a_Y^*(\lambda)$ during inverse modeling of $E_d(\lambda)$ measurements.

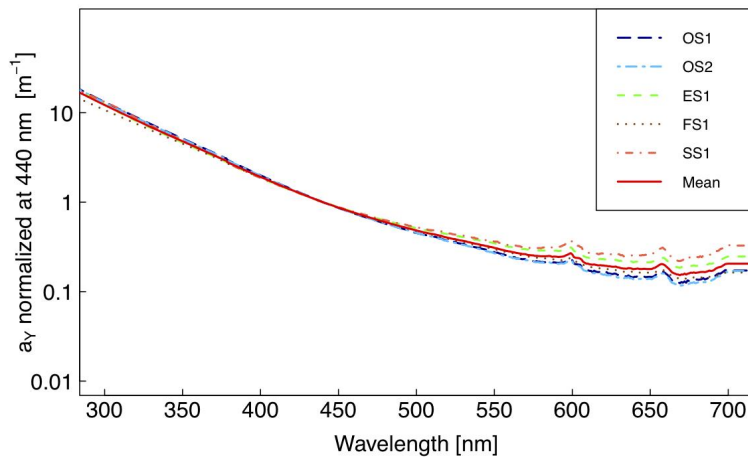


Figure 4. Mean specific absorption coefficient $a_Y^*(\lambda)$ of each sampling site normalized at $\lambda_0 = 440 \text{ nm}$. The red line represents the mean specific absorption coefficient $a_{Y,OS}^*(\lambda)$ of all CDOM measurements ($N=60$).

4.2 Validation of fit results

The results of laboratory analysis of the water samples, which were presented in section 4.1, are used to validate the corresponding parameters obtained by inverse modeling of downwelling irradiance measurements. Fit results from E_d measurements performed at a water depth of 1 m or less showed strong fluctuation of the fitted $a_Y(440)$. Therefore only fit results from 1.5 m water depth or more are considered, leaving $N = 24$ samples for the comparison.

In order to account for possible sensitivity of the fit results on initialization, data analysis was conducted for different initial values for Y and C_i , while the initial values for $X, f_{dd}, f_{ds}, \alpha, \beta, H_{oz}, WV$ and z were not changed. The influence of the specific absorption coefficient of CDOM, $a_Y^*(\lambda)$, was investigated by using different S values in the exponential approximation, treating S as fit parameter, and using the measured spectrum $a_{Y,OS}^*(\lambda)$ from Figure 4 instead of the exponential function. In Table 2 the fit results of the different settings are summarized.

Table 2. Correspondence between $a_Y(440)$ obtained from inverse modeling of E_d measurements and laboratory analysis of water samples. The columns represent different settings of the inversion algorithm. In the standard setting, CDOM concentration was initialized by $Y = 0.5 \text{ m}^{-1}$, the phytoplankton concentrations by $C_i = (1, 0, 0, 0) \text{ mg m}^{-3}$ ($i = 0, 3, 4, 5$), and $a_Y^*(\lambda)$ was calculated as $\exp[-S(\lambda-440)]$ with constant $S = 0.0140 \text{ nm}^{-1}$.

Changed fit parameter	Different settings											
	Y initial value [m^{-1}]					S set value [nm^{-1}]					C _i classes	
	0.5'	0.002	1	5	10	0.014'	0.004	0.025	fitted	$a_{Y,OS}^*$	C [0]'	C [3,4,5]
N	24	24	24	24	24	24	24	24	24	24	24	24
coefficient of determination	0.92	0.00	0.90	0.90	0.00	0.92	0.64	0.80	0.91	0.87	0.92	0.93
slope	0.80	-0.02	0.81	0.80	0.11	0.80	0.49	1.38	0.76	0.91	0.80	0.74
y-intercept [m^{-1}]	0.08	0.20	0.09	0.08	1.27	0.08	0.30	0.34	0.13	0.01	0.08	0.12
standard deviation of fitted $a_Y(440)$ [m^{-1}]	0.06	0.14	0.05	0.07	1.29	0.06	0.11	0.13	0.06	0.06	0.06	0.06

standard setting ' 1

The standard setting yields a high coefficient of determination of $R^2 = 0.92$, a slope of 0.80, a y-intercept of 0.08 m^{-1} , and a standard deviation of 0.06 m^{-1} (Figure 5a). Changing Y initial values from the standard value of 0.5 m^{-1} to the extreme values of 0.002 m^{-1} and 10 m^{-1} destroys the correlation. Y initial values of 1 m^{-1} and 5 m^{-1} affect the fit results only slightly, i.e. the inversion algorithm is quite robust concerning Y initialization and tolerates deviations by an order of magnitude from the real concentration. Unrealistic slope values decrease the correlation and increase the standard deviation of the fit results. Fitting the slope S in addition to the other fit parameters results in no better correlation. The best regression results are obtained by taking the mean specific absorption spectrum of the Osterseen Lake District, $a_{Y,OS}^*(\lambda)$ (Figure 5c). The slope of the regression line is closer to 1 than for any other setting (0.91), and the y-intercept is very close to zero (0.01 m^{-1}). Treating the three typical phytoplankton classes separately declines the regression parameters, indicating that no further differentiation between the phytoplankton classes can be performed using the applied settings of the inversion algorithm.

The correlation result from the standard setting and using $a_{Y,OS}^*(\lambda)$ is illustrated in Figure 5a and 5c. Different sampling depths are marked in different colors. Additionally the dependency of the standard deviation of the fitted $a_Y(440)$ values on sensor depth is shown in Figure 5b and 5d. Gege¹⁴ already observed that a path length of 1 to 1.5 m is required to derive reliably absorption from E_d measurements. This result can be confirmed for the CDOM retrieval of this study.

A similar analysis was made for S . Fitted S values and S values obtained from water samples are only weakly correlated ($R^2 = 0.23$, slope = 0.143, y-intercept = 0.01 nm^{-1}), i.e. S cannot be determined reliably by inverse modeling of E_d measurements.

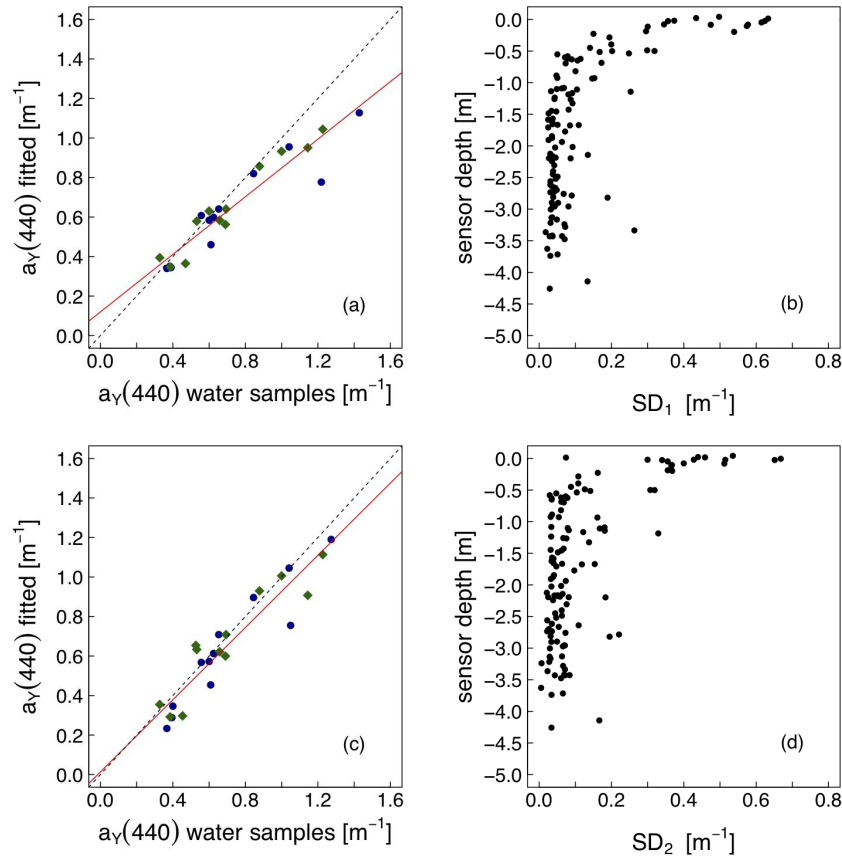


Figure 5. Illustration of the two best fit results. (a) $a_V(440)$ determined using the standard setting. (b) The calculated standard deviation of the fitted CDOM concentrations as function of sensor depth. (c) $a_V(440)$ determined using $a_{Y,OS}^*(\lambda)$ and (d) the corresponding standard deviation as function of sensor depth. Different colors and symbols represent the two sampling depths, 2 m (blue points) and Secchi disk depth (green rhombs).

5. CONCLUSION

The potential of a new analytical model for downwelling irradiance in water, which can cope with the large and unpredictable fluctuations of the underwater light field, for estimating CDOM was analyzed in this study. The accuracy of the derived parameters was determined by comparison with independent laboratory measurements of water samples, which were acquired simultaneously to the underwater spectral measurements. The campaigns were performed in different lakes of the Osterseen Lake District in summer 2012. Realistic concentrations were obtained for sampling depths above 1.5 m for all sampling sites. The accuracy depends on the knowledge of the spectral shape of CDOM absorption, represented by the spectrum $a_Y^*(\lambda)$. The standard exponential approximation with the spectral slope $S = 0.0140 \text{ nm}^{-1}$ provided already good results ($R^2 = 0.92$) because this S value is close to the average of 0.0145 nm^{-1} of the sampling lakes. Replacing the exponential $a_Y^*(\lambda)$ model by the average spectrum $a_{Y,OS}^*(\lambda)$ for the Osterseen Lake District clearly improves the fit results. The spectrum $a_{Y,OS}^*(\lambda)$, which has been determined in this study, differs slightly from an exponential function. Interestingly, changing the phytoplankton classes did not improve the fit results. Hence the default spectrum of phytoplankton absorption in WASI, which is a weighted sum of five phytoplankton spectra from Lake Constance, can also be used for the freshwater lakes studied here. The spectral slope S could not be determined accurately by inverse modeling of the irradiance spectra, but as correlation with the laboratory results was significant, it may be worth to spend further investigation on this matter in future. Overall this new model of downwelling irradiance can be used as an alternative to laboratory analysis of water samples to determine CDOM concentration.

6. ACKNOWLEDGMENT

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