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2 **Effects of dams on riverine biogeochemical cycling and ecology**

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30 **Abstract**

31 Currently, dam construction is the main and growing global anthropogenic disturbance on
32 rivers. Dams have major effects on the physics, chemistry and biology of the original river
33 including altering water circulation and retention time, sedimentation, nutrient
34 biogeochemical cycling including greenhouse gas emissions, and the amount and
35 composition of the organisms. Among those, the effect of dams on riverine material cycle
36 and ecology is especially concerning because of its close relationship with current global
37 environmental problems such as climate change and ecological deterioration. This review
38 thus mainly focuses on nutrient cycling and ecological changes in a regulated river. In the
39 future, researches on a reservoir-river system should focus on: (1) processes and
40 mechanisms of nutrient biogeochemical cycles; (2) interaction between these processes and
41 ecological change such as phytoplankton succession; (3) developing mathematical
42 functions and models to describe and forecast these processes and their interaction in the
43 future.

44 **Key words:** dam; retention time; nutrient; greenhouse gas; ecosystem

45

46 **1. Introduction**

47 Rivers are the major links connecting the land to the ocean; they deliver fresh water, carbon,
48 energy and nutrients to the estuary and coastal sea (e.g., [Humborg et al., 1997](#); [Jiao et al., 2007](#)). Rivers also link the land with the atmosphere, exchanging heat and influencing the
49 regional climate, and exchanging gases, affecting global biogeochemical cycles and the
50 global climate ([Lauerwald 2015](#)). In the past decades, with the increasing demands of a
51 growing human population, the natural river has been strongly disturbed by dam
52 construction to generate hydropower, increase water supply and security, control floods,
53 improve navigation and provide opportunities for recreation ([Bednarek 2001](#)). Thus, the
54 natural connectance between the land, rivers and the ocean has declined and material
55 cycling has been affected with important consequence for the biology of the altered
56 ecosystems.

58 However, until the 1970s, the environmental impacts of dams were not widely taken
59 into account. In 1972, the Scientific Committee on Problems of the Environment issued a
60 report of man-made lakes as modified ecosystems ([SCOPE Working Group on Man-made
61 Lakes, 1972](#)), which showed earlier concern about the physical, chemical, and biological
62 impacts of dams on the downstream rivers. In 2000, the World Commission on Dams

63 presented another report about dams and development: a framework for decision making
64 ([The report of the world commission on dams, 2000](#)). In 2010, the International
65 Hydropower Association proposed the hydropower sustainability assessment protocol.
66 Gradually, the effect of dams on rivers and their connected ecosystems becomes a popular
67 scientific topic ([IHA, 2010](#)). The Millennium Ecosystem Assessment noted the dramatic
68 increase in dam construction and consequent water storage to the extent that flows in 60%
69 of the World's large rivers are moderately or strongly affected and also noted some of the
70 negative ecosystem consequences that this has caused ([Millennium Ecosystem Assessment](#)
71 [2005](#)).

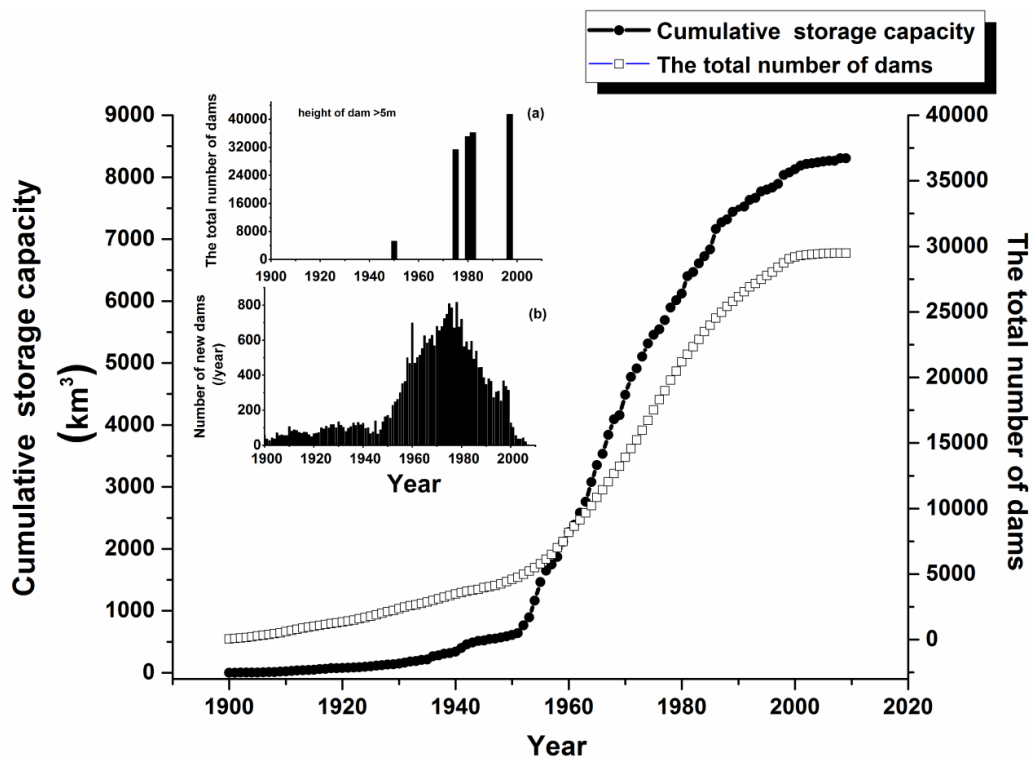
72 Generally, dam construction has three consequences: i) altered river hydrological cycle,
73 exacerbated by artificial regulation such as anti-seasonal storage; ii) altered biogeochemical
74 cycles in the impounded river; iii) altered ecological conditions in the discontinuous river-
75 reservoir system. These processes interact with each other, and their influences can have
76 local, regional and global effects. The understanding of these processes is the scientific
77 basis for explaining the environmental impacts of dam construction and producing
78 sustainable management strategies for the impounded river. This brief review mainly
79 focuses on these processes and, in addition, current hot topics about the impounded river
80 are also discussed.

81 **2. Historical and current states of river damming**

82 Modern dam construction began in 1900 and boomed from about 1950 with the use of
83 concrete and innovation of excavation ([Fig.1](#)). Currently, approximately 70% of the world's
84 rivers are intercepted by dams ([Kummu and Varis 2007](#)), and in China, there were more
85 than 80,000 reservoirs by the end of 2008, among which there were over 5000 dams with
86 a height of over 30 m (<http://www.chincold.org.cn>). Dams are built to store water for
87 various purposes. Accompanied with the rapid increase of dam construction (from 1948 to
88 2010), the global active storage capacity of reservoir grew from about 200 to over 5000
89 km³, which is over 70% of the total global reservoir capacity (7000-8000 km³) ([Vörösmarty](#)
90 [1997](#); [Zhou et al., 2016](#)). The number of reservoirs will increase in the future with the restart
91 of hydropower loan project by the World Bank ([World Bank 2009](#)) and the motivation to
92 increase renewable energy sources ([Hermoso 2017](#)).

93 Globally, the extent of hydropower development is not balanced. In Europe, North
94 America and Central America, more than 70% of the technically feasible hydropower has
95 been utilized, while this value is less than 4% in Africa ([Wang and Dong 2003](#), [Home 2005](#)).
96 The developed countries have higher level of hydropower utilization than the developing

97 countries. For example, in China, only about 24% of the hydropower resource has been
 98 exploited, much less than the average value of 60% in the developed countries. There are
 99 also large regional differences within China: eastern China has exploited 79.6% of its
 100 hydropower resources, but southwestern China, which has the richest hydropower
 101 resources, has only exploited 8.5% (Liu et al., 2009). With the development of the global
 102 economy, especially in the developing countries, it can be expected that the rivers will face
 103 large-scale dam construction, and human disturbance will be further intensified on the
 104 rivers (Zarfl et al., 2015; Hermoso 2017).



105
 106 **Figure. 1.** Historical variation in the number and cumulative storage capacity of reservoirs
 107 (modified from (Chao et al., 2008; Jia et al., 2010). (a) Number of dams in the world with a dam
 108 height > 5 m (Liu et al., 2009; Jia et al., 2010); (b) The number of new dams constructed in the
 109 world each year and (c) The total number of dams and the cumulative storage capacity of reservoirs
 110 registered in the ICOLD (Chao et al., 2008).

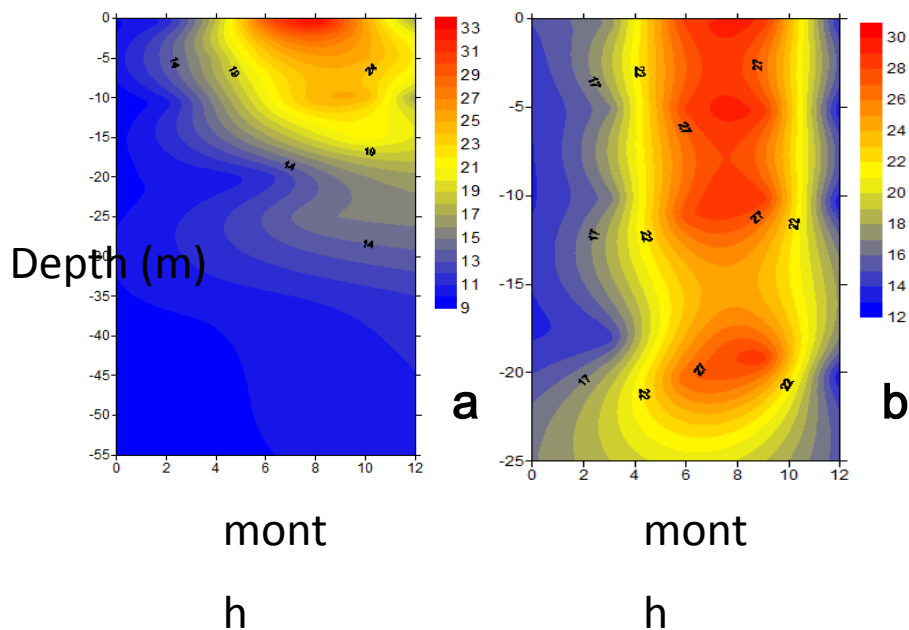
111 **3. Impacts of river damming on the hydrological cycle and physical characteristics**

112 **3.1 Seasonal thermal stratification**

113 Reservoir stratification conforms to the classical pattern of lake stratification especially
 114 hydroelectric reservoirs that are usually deep and thus usually develop seasonal thermal
 115 stratification. The densimetric Froude Number (F) has been suggested to estimate the
 116 stratification tendencies in a reservoir (Ledec and Quintero 2003). The stratification is

117 expected when F is less than 1, the severity of which increases with a smaller F ; while when
 118 F is greater than 1, stratification is not likely. For hydroelectric reservoirs, the extent of
 119 thermal stratification is influenced by the pattern and extent of water storage and discharge
 120 (Fig. 2). One consequence of thermal stratification is that if water is released from the
 121 bottom of the reservoir during stratification it will be very different to that at the surface
 122 with potential effects on the downstream river for tens of kilometers (Petts 1984). It has
 123 been suggested that this problem could be eliminated by artificially destroying thermal
 124 stratification (Lackey 1972; Elçi2008), or by releasing water from the surface or sub-surface.

125 A reservoir can also have an effect on the downstream river temperature with
 126 consequences for biogeochemical cycling and river ecology and particularly fish
 127 populations. In general, water released from the bottom of a reservoir will be cooler than it
 128 would be without the reservoir while water released from the surface of the reservoir will
 129 be warmer. However, in addition, reservoirs also dampen the temperature cycle at seasonal,
 130 daily and sub-daily timescales and, in a study of Canadian reservoirs, increased the mean
 131 water temperature in September (Maheu et al., 2016).

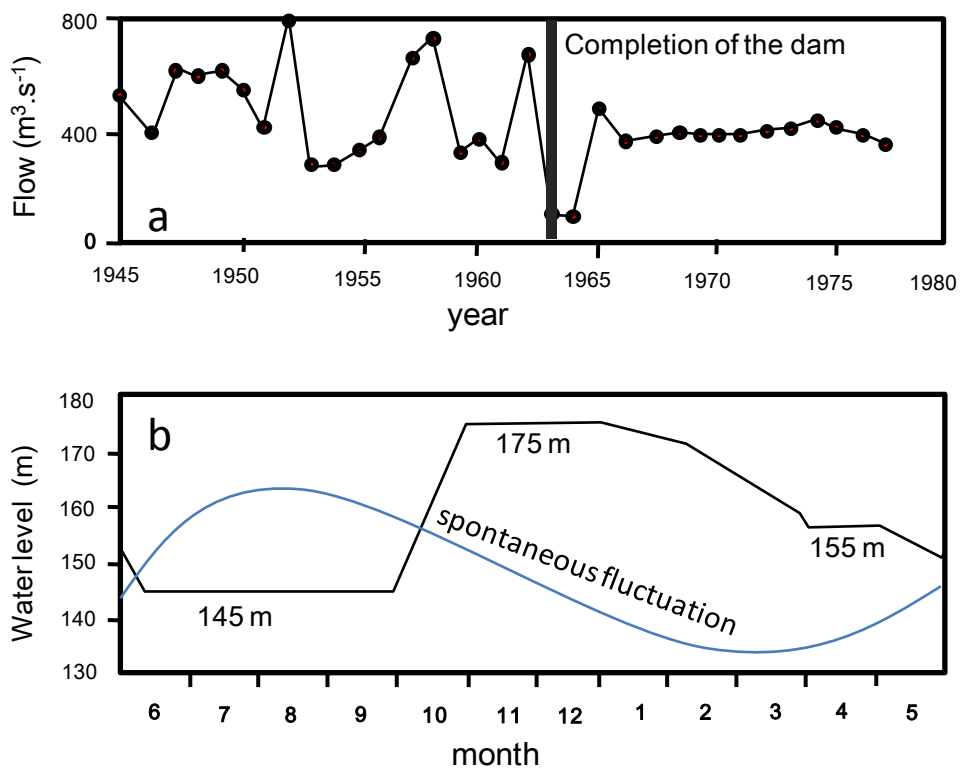


132
 133 **Figure. 2.** Thermal stratification in different reservoirs (unpublished data). (a) The Xinanjiang
 134 Reservoir with a water retention time greater than one year; (b) The Wanan Reservoir with a
 135 water retention time less than two weeks.

136 3.2 Storage pattern

137 Artificial regulation of reservoir water, such as storing or releasing water, changes the flood
 138 pulse of the original river, affects the water balance of the basin and the hydrological

139 condition of the river bank (Liu et al., 2009) . After interception by a dam, a significant
 140 reduction in the maximum flow is found in the downstream river (Fig. 3a). In addition,
 141 large and medium-sized reservoirs often have anti-seasonal storage to reduce reservoir
 142 water level in flood season to cope with flood peaks (e.g., Fig. 3b). This is totally different
 143 from a lake; where water level changes corresponding to the input or river water minus
 144 evaporation. When cascade reservoirs are constructed, competitive water storage among
 145 the reservoirs will occur. This will enhance the disadvantage of requirement of water in the
 146 downstream reaches of the river.



147
 148 **Figure 3.** Changes in river flow and water level after river damming. (a) Effects of an upstream reservoir
 149 (Lake Powell) on the maximum annual flow of downstream (Modified from Thornton et al., 1996). (b)
 150 Water level change after the Three Gorge Reservoir closure.

151 3.3 Hydrological retention time

152 Dam construction obviously changes the retention time of the corresponding river. Within
 153 one reservoir, the retention time can vary between one day and several years, greatly
 154 prolonging that of the natural river. For continental runoff in free-running river channels,
 155 the average residence time varies from <16 to 26 days, while the discharge weighted
 156 global average value is almost 60 days (Vörösmarty and Sahagian 2000). In some
 157 strongly regulated river basins, the value can be higher. For example, the water retention
 158 time of the Huanghe River (upstream of Lijin station, i.e. the whole basin taking into

159 account the 2816 reservoirs), increased from one year to 4 years after dam construction,
 160 ranking the Yellow River in the top three in terms of residence time and flow regulation
 161 among large river systems in the world (Ran and Lu 2012). The increase of the water
 162 retention time has a profound effect on the reservoir thermal stratification, riverine
 163 elemental cycle, and phytoplankton ecology (e.g., Duras and Hejzlar 2001; Soares et al.,
 164 2012).

165 **4. Effect of dams on riverine material cycling**

166 **4.1 Sedimentation**

167 Sedimentation within reservoirs is a complex process. The sediment load delivered to the
 168 reservoir is controlled by the sediment yield in a basin, and reservoir sedimentation is
 169 mainly influenced by the hydraulics of the river, the geometry of the reservoir and ratio at
 170 the entrance to reservoir (width to depth ratio). The distribution of sediment types in a
 171 reservoir is shown in Table 1. The reduction in downstream sediment load by reservoir
 172 construction may be greater than 75%, as seen in the case studies of Sao Francisco River
 173 in Brazil, the Chao Phraya River in Thailand, and the Yellow River in China (Walling 2006).
 174 Kumm and Varis (2007) reported that the operation of dams on the Mekhong main channel
 175 had approximately halved the sedimentation from $150\text{-}170 \times 10^9$ to 81×10^9 kg annually. The
 176 magnitude of global suspended sediment flux to the ocean is still unclear and has been
 177 estimated to be in the range of 9.3 Gtyr^{-1} (Judson 1968), to more than 58 Gtyr^{-1}
 178 (Holeman 1968) with recent studies converging around $15\text{-}20 \text{ Gtyr}^{-1}$ (e.g., Milliman and
 179 Meade 1983; Meybeck 1988; Ludwig et al., 1996; Vörösmarty et al., 2003).

180

181 **Table 1** A typical distribution of deposited sediment in a reservoir (USACE 1987).

Particle size	Inlet (%)	Mid-reservoir (%)	Outlet (%)
Sand	5	<1	0
Silt	76	61	51
clay	19	38	49

182

183 **4.2 Nutrient retention**

184 The geochemical behaviors of nutrient elements such carbon (C), nitrogen (N), silicon (Si),
 185 and phosphorus (P) are obviously influenced by dam construction. After damming a river,
 186 particulate nutrients are partially impounded and their fluxes through the river thus decrease.

187 In addition, biological activity such as photosynthesis and respiration modifies the
188 dissolved nutrient content and species in the reservoir. These modified nutrients are then
189 transported to the downstream river, and finally influence the ecosystem of the estuary and
190 marginal sea (e.g., [Humborg et al., 1997](#); [Jiao et al., 2007](#)).

191 The first concern on the consequences of dams on nutrient retention was about Si. Dam
192 construction resulted in a great deal of Si retention and a significant decrease in the flux of
193 dissolved Si (DSi) to the sea ([Humborg et al., 1997, 2006](#); [Wang et al., 2010](#); [Maavara et al., 2014](#)).
194 At the global scale, the retention of DSi in lakes and reservoirs is 163 Gmol yr^{-1}
195 ($9.8 \text{ Tg SiO}_2 \text{ yr}^{-1}$), and the total active Si retained is 372 Gmol yr^{-1} ($22.3 \text{ Tg SiO}_2 \text{ yr}^{-1}$)
196 ([Maavara et al., 2014](#)).

197 Recently, [Maavara et al. \(2015, 2017\)](#) estimated the global P and organic C (OC)
198 retention by river damming. Total P (TP) trapped in the global reservoirs was estimated to
199 be 22 Gmol yr^{-1} in 1970 and 42 Gmol yr^{-1} in 2000, and active P was 9 Gmol yr^{-1} in 1970 and
200 18 Gmol yr^{-1} in 2000; however, the global river TP load to the ocean in the same period,
201 only changed from 312 to 349 Gmol yr^{-1} ([Maavara et al., 2015](#)). The rapid increase of total
202 and active P retention was mainly caused by the rapid expansion of dam construction
203 between 1970 and 2000, and the volume of reservoirs increased from about $3,000$ in 1970
204 to almost $6,000 \text{ km}^3$ in 2000 ([Lehner et al., 2011](#)). By 2030, about 17% of the global river
205 TP load will be sequestered in reservoir sediments, and the main increase is from Asia and
206 South America, especially in the Yangtze, Mekong, and Amazon drainage basins ([Maavara et al., 2015](#)).
207 As for global OC, its mineralization in reservoirs exceeds C fixation, and
208 about 75% of OC in reservoir sediments is allochthonous. OC burial in reservoirs is
209 estimated to be about 4.3 Tmol yr^{-1} by 2030, a fourfold increase relative to 1970, and OC
210 mineralization is around 2.6 Tmol yr^{-1} in the same time. The total value (6.9 Tmol yr^{-1})
211 accounts for about 19% of total OC carried by rivers to the oceans by 2030 ([Maavara et al., 2017](#)).
212 Comparatively little is known about N retention in impounded rivers, perhaps
213 because of the more complex N cycle in the reservoir-river system. A case study in a
214 regulated Mediterranean river indicated that river reaches below dams acted as net sinks of
215 total dissolved N, unlike dissolved P or OC, and this high net uptake by organisms
216 (autotrophs and heterotrophs) below dams could reduce N export to downstream
217 ecosystems ([von Schiller et al., 2016](#)).

218 4.3 Greenhouse gas emissions

219 Reservoir greenhouse gases (GHGs) based on carbon, CO_2 and CH_4 , are derived from OC
220 mineralized in the reservoir or direct input of CO_2 produced in the catchment ([Maberly et](#)

221 al., 2013), and their emission occurs by diffusion across the air-water interface and,
222 especially for CH₄, ebullition. GHG production and emission fluxes from a reservoir are
223 closely related to reservoir age, latitude, and retention time (Barros et al., 2011; Ometto et
224 al., 2013; Wang et al., 2015; Deemer et al., 2016). The reservoir surface is usually
225 dominated by the diffusive flux of CO₂, even in cases in which bottom anoxia leads to high
226 CH₄ production because of conversion of upwardly-diffusing CH₄ to CO₂ by
227 methanotrophic bacteria. However, when water is released from the bottom of the dam,
228 CH₄ emissions can be very high. It is suggested that this source contributes 50 to 90% of
229 total CH₄ emissions from tropical or temperate hydroelectric reservoirs (Abril et al., 2005;
230 Kemenes et al., 2007; Maeck et al., 2013).

231 Rates of surface diffusion of GHGs among reservoirs have been shown to vary along
232 broad geographic gradients, low-latitude tropical reservoirs typically emitting GHGs at
233 greater rates per unit area than high-latitude temperate and boreal reservoirs (Barros et al.,
234 2011). Average emissions of 3500 mg m⁻² d⁻¹ of CO₂ and 300 mg m⁻² d⁻¹ of CH₄, have been
235 found in tropical reservoirs compared to 387~1400 mg CO₂ m⁻² d⁻¹ and 2.8~55 mg CH₄ m⁻²
236 d⁻¹ from temperate reservoirs (mostly hydroelectric reservoirs) (St. Louis et al., 2000;
237 Soumis et al., 2005; Lima et al., 2008; Barros et al., 2011). Chanudet et al. (2011) estimated
238 diffusive fluxes to the atmosphere from two Southeast Asian sub-tropical reservoirs to be -
239 466~1680 mg m⁻² d⁻¹ for CO₂ and 12.8~190 mg m⁻² d⁻¹ for CH₄, comparable to other
240 tropical reservoirs. Few studies have focused on reservoir N₂O emission. A case study in
241 the Wujiang cascade reservoirs showed that the average flux of N₂O emission from the
242 reservoir surface was about 0.64 μmol m⁻² h⁻¹, comparatively to the natural lake (Liu et al.,
243 2011).

244 It has been estimated that hydropower reservoirs account for 30%- 62% of all
245 reservoirs globally (Lehner et al., 2011; Varis et al., 2012). Lima et al. (2008) first estimated
246 that CH₄ emission from the global hydroelectric reservoirs are 100 Tg CH₄-C yr⁻¹; however,
247 Barros et al. (2011) suggested this value was 3 Tg CH₄-C yr⁻¹, with CO₂ emission of 48 Tg
248 CO₂-C yr⁻¹ in the hydroelectric reservoirs. While the values from Herwich (2013) were 76
249 Tg CO₂-C yr⁻¹ and 7.3 Tg CH₄-C yr⁻¹, respectively. The large range in published estimates
250 could be caused by the different estimate of global reservoir surface (Mendonça et al., 2012;
251 Teodoru et al., 2012; Mosher et al., 2015). Perhaps high heterogeneity of CO₂ and CH₄
252 fluxes along the long and narrow reservoirs also contributes to these differences.

253 **5. Effect of dams on riverine ecology**

254 **5.1 River continuum concept**

255 The river continuum concept (RCC) describes the longitudinal gradient of physical
256 conditions such as geomorphological and hydrological factors in pristine rivers (Vannote
257 et al., 1980; Tornwall et al., 2015). Biological communities are adapted to these gradients,
258 and vary along the river from the headwaters to the mouth in a predictable manner. The
259 headwater regime is strongly heterotrophic (the ratio of photosynthesis to respiration (P/R)
260 <1), and has coarse particular matter and invertivores as the main biological species. The
261 mid-regime is autotrophic (P/R>1), and has fine particular matter and piscivorous,
262 invertivorous, and planktivorous species. Finally, the downstream regime gradually return
263 to heterotrophy due to turbidity (Fisher 1977; Vannote et al., 1980). Dams interrupt the river
264 continuum altering geomorphology, water quality, temperature regime, and flow regime,
265 and result in upstream–downstream shifts in biotic and abiotic patterns and processes. The
266 serial discontinuity concept (SDC) views impoundments as major disruptions to
267 longitudinal resource gradients along river courses (Ward and Stanford 1983, 1995). The
268 impacts under impoundment have been studied with respect to geomorphology,
269 temperature, flow, invertebrates, and fish (e.g. Kondolf 1997; Jakob et al., 2003; Poff and
270 Zimmerman 2010; Jones 2011; Winemiller et al., 2017). Usually, periphyton biomass
271 recovers quickly within 5 km downstream, while benthic invertebrate richness varies
272 considerably, with both increases and reductions observed at near-dam sites and varying in
273 recovery downstream (Ellis and Jones 2013, 2016).

274 **5.2 Phytoplankton**

275 As the main primary producers, river phytoplankton succession after damming is an
276 important issue. Dams can result in major changes in the phytoplankton community in the
277 river, estuary and adjacent sea (e.g., Humborg et al., 1997; Jiao et al., 2007). In a river-
278 reservoir unit, the dominant Bacillariophyta (diatoms) in river water changes to co-existing
279 Bacillariophyta, Chlorophyta (green algae), and Cyanophyta (blue-green algae) in
280 mesotrophic reservoirs, and then shifts to dominance by Cyanophyta in eutrophic reservoirs
281 (Wang et al., 2013). Phytoplankton succession in the impounded river is not directly caused
282 by the physical obstruction of the dam, but can be attributed to changes in hydrological and
283 geochemical conditions after damming. For example, phytoplankton community
284 succession in karst cascading reservoirs was influenced by Si and P stoichiometry (Wang
285 et al., 2014a), while in a tributary of the Three Gorges Reservoir, phytoplankton diversity
286 was controlled by hydraulic retention time and nutrient limitation (Xiao et al., 2016).
287 Phytoplankton dynamics in the impounded river is non-linear, and the mechanisms
288 responsible need further research. Generally, the study of the effects of dams on riverine
289 ecology is still at an initial stage, especially in the aspect of the coupling of ecological shifts

290 and nutrient biogeochemical cycle.

291 **5.3 Fish**

292 The effect of dams on fish ecology (e.g. spawning, migration, and diversity) has always
293 been a concern. Globally, dam obstruction decreases fish biodiversity, and taxa such as
294 lampreys (*Lampetra* spp.), eels (*Anguilla* spp.), and shads (*Alosa* spp.) are at particular risk
295 of species loss (Fu et al., 2003; Liermann et al., 2012). Regionally, the construction of the
296 Three Gorges Reservoir, for example, resulted in a substantial decline in carp larval
297 abundance of the middle Yangtze River (Wang et al., 2014b).

298 **6. Perspectives**

299 The construction of dams for various purposes, but particularly hydropower, is booming
300 (Zarfl et al., 2015) and is likely to accelerate further. Given the complex interactions
301 between land, rivers, the estuary and the atmosphere, the consequences of dam building
302 will inevitably have complex knock-on, ecological and social effects. International
303 initiatives such as the Paris Agreement reached at the COP21 in December 2015 are
304 encouraging countries to move towards a greater reliance on renewable energy production.
305 Hydropower currently accounts for more than 80% of renewable energy (Zarfl et al., 2015).
306 Hermoso (2017) pointed out that while this might prove beneficial for global carbon
307 emissions it would be likely to prove detrimental to local freshwater ecosystems and
308 consequently there is a requirement for international guidance and legislation in order to
309 evaluate the benefits of construction of a new dam compared to the ecological and societal
310 costs. For example, proposals to increase dam construction on the Amazon River led
311 Latrubesse et al. (2017) to introduce a ‘Dam Environmental Vulnerability Index’ based on:
312 (1) the vulnerability of the basin to run-off and erosion that may transport nutrients and
313 pollutants to the river, (2) modification to the hydrological regime and transport of sediment,
314 (3) quantification of the extent of the river system affected. Clearly, there is a growing and
315 urgent need for robust scientific evidence to make this kind of evaluation. Currently, while
316 numerous case-studies exist at specific sites, it is difficult to take account of regional
317 heterogeneity in conditions in order to produce advice at a global scale. In the area of the
318 role of reservoirs in global biogeochemical cycling, future research should focus on the
319 following aspects: (1) processes and mechanisms of nutrient biogeochemical cycles; (2)
320 coupling these cycles with ecological conditions such as phytoplankton succession; (3)
321 developing mathematical functions and models to describe and forecast these processes and
322 their interaction in the future

323

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329

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