

**Faculty of Science and Engineering
Department of Petroleum Engineering**

**Pore-Scale Investigations of Rock and Fluid Microstructure and Fluid
Displacement Processes in Geological Porous Media**

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**This thesis is presented for the Degree of
Doctor of Philosophy
of
Curtin University**

August 2018

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DEDICATION

I would like to dedicate my thesis to my beloved family.

To my father, Mr. Anwar Abubakar BA. He is the best role model, and the greatest inspirator for my whole entire life.

To the people of Aceh who have sacrificed their lives for freedom, independence, sovereignty, justice, and dignity of their country!

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ABSTRACT

Since the beginning of the industrial age, excessive use of fossil fuels has led to a significant increase in the atmospheric concentration of greenhouse gases. Sequestration of carbon dioxide in geological formations appears to be the most direct carbon-management strategy for reducing such anthropogenic CO₂ in the long term. Residual trapping mechanisms have been recognised as effective approaches for CO₂ geo-sequestration projects, as they immobilise micrometre-sized CO₂ bubbles in the pore network of the rock and increase the safety of the storage sites as they inhibit upward migration of CO₂, thus, minimising leakage potential. This mechanism has been proven to work in water-wet rocks with no existing works. However, no such work has been completed for (oil-wet) oil reservoirs, which are economically most important, and it was hypothesised that this significantly affects residual trapping.

Thus, this study evaluated the supercritical carbon dioxide trapping capacity of oil-wet formations at reservoir conditions via x-ray micro-computed tomography. The CO₂ clusters and their morphologies were analysed statistically and it was clear that residual saturations are approximately halved in oil-wet rock while cluster morphology changes into flatter clusters.

These findings need to be implemented in reservoir-scale models to obtain accurate results.

List of Publications included as part of the thesis

1. Al-Yaseri, A.Z., Roshan, H., Zhang, Y., **Rahman, T.**, Lebedev, M., Barifcani, A. and Iglauer, S., 2017. Effect of the Temperature on CO₂/Brine/Dolomite Wettability: Hydrophilic versus Hydrophobic Surfaces. *Energy & Fuels*, 31(6), pp.6329-6333.
2. Lebedev, M., Zhang, Y., Mikhaltsevitch, V., Iglauer, S. and **Rahman, T.**, 2017. Residual trapping of supercritical CO₂: direct pore-scale observation using a low-cost pressure cell for micro-computer tomography. *Energy Procedia*, 114, pp.4967-4974.
3. **Rahman, T.**, Lebedev, M., Barifcani, A. and Iglauer, S., 2016. Residual trapping of supercritical CO₂ in oil-wet sandstone. *Journal of Colloid and Interface Science*, 469, pp.63-68.
4. Iglauer, S., **Rahman, T.**, Sarmadivaleh, M., Al-Hinai, A., Fernø, M.A. and Lebedev, M., 2016. Influence of wettability on residual gas trapping and enhanced oil recovery in three-phase flow: a pore-scale analysis by use of microcomputed tomography. *SPE Journal*, 21(06), pp.1-916.

List of Publications relevant to the thesis but not forming part of it

1. **Rahman, T.**, Lebedev, M., Zhang, Y., Barifcani, A. and Iglauer, S., 2017. Influence of rock microstructure on its electrical properties: an analysis using x-ray microcomputed tomography. *Energy Procedia*, 114, pp.5023-5031.
2. Zhang, Y., Lebedev, M., Sarmadivaleh, M., Barifcani, A., **Rahman, T.**, and Iglauer, S., 2016. Swelling effect on coal microstructure and associated permeability reduction. *Fuel*, 182, pp.568-576.

TABLE OF CONTENTS

DECLARATION OF ACADEMIC INTEGRITY	I
COPYRIGHT	II
DEDICATION.....	III
ACKNOWLEDGMENT	IV
ABSTRACT	V
TABLE OF CONTENTS.....	VIII
List of Figures.....	XI
List of Tables	XVII
Chapter 1 Introduction.....	1
1.1 Background	1
1.2 Research objective.....	12
1.3 Thesis organisation.....	12
Chapter 2 Literature Review	15
2.1 Residual Trapping of Supercritical CO ₂	15
2.1.1 Storage Mechanisms	20
2.2 Influence of Wettability on Residual Gas Trapping and EOR.....	23
2.3 Dolomite Wettability after CO ₂ injection.....	27
Chapter 3 Effect of Temperature on CO₂/Brine/Dolomite Wettability: Hydrophilic vs Hydrophobic Surfaces*	29
3.1 Introduction	30
3.2 Experimental methodology	31
3.3 Theoretical background.....	34
3.4 Results and discussion.....	36
3.5 Conclusions	40

Chapter 4 Residual trapping of supercritical CO₂: direct pore-scale observation using a low-cost pressure cell for micro-computer tomography*	41
4.1 Introduction	42
4.2 Pressure cells for microCT	43
4.3 Demonstration experiments	46
4.4 Conclusion	50
Chapter 5 Residual Trapping of Supercritical CO₂ in Oil-wet Sandstone*	51
5.1 Introduction	52
5.2 Experimental Procedure	53
5.2.1 Simulating Subsurface Condition	53
5.2.2 Imaging scCO ₂ at the pore-scale with μ CT	53
5.3 Results and Discussion	54
5.3.1 Visualization, cluster morphologies and initial & residual CO ₂ Saturations	54
5.3.2 CO ₂ cluster statistics	58
5.4 Conclusions and Implications	62
Chapter 6 Influence of wettability on residual gas trapping and enhanced oil recovery in three-phase flow: a pore-scale analysis using micro-computed tomography*	63
6.2 Experimental Methodology	65
6.2.1 Waterflood	66
6.2.2 Gasflood	66
6.3 Results and Discussion	67
6.3.1 Qualitative analysis Pore-scale fluid configurations	72
6.3.2 Quantitative analysis - Cluster statistics	79
6.3.3 Interface curvatures and capillary pressures	84
6.4. Conclusions	87
Chapter 7 Conclusion and Future Work	89

7.1	Conclusion.....	89
7.2	Future Work	92
	References	93
	Appendices	118
	Appendix 1: Statement of Contributions of Others	118
	Appendix 1.1: Statement of Contributions of Others for	119
	Appendix 1.2: Statement of Contributions of Others for	125
	Appendix 1.3: Statement of Contributions of Others for	130
	Appendix 1.4: Statement of Contributions of Others for	134
	Appendix 2: Copyright Forms.....	140
	Appendix 2.1.....	141
	Appendix 2.2.....	142
	Appendix 2.3.....	143
	Appendix 2.4.....	144

List of Figures

- Figure 1-1** Images of CO₂ after waterflooding with CO₂ saturated brine, showing the trapped supercritical CO₂ clusters. The picture is a two-dimensional cross-section through a three-dimensional image; the area shown is 10.80 mm² (261 239 voxels). (a) A raw grey-scale image: supercritical CO₂ is black, brine is light grey, and sandstone is dark grey. The few white areas are minerals with high X-ray absorption. (b) The same slice with phases segmented. Supercritical CO₂ is white, brine light blue, and rock is brown (figure taken from Blunt et al. (2013)). 3
- Figure 1-2** Images of CO₂ after waterflooding with CO₂ saturated brine showing selected trapped clusters of different sizes (figure taken from Blunt et al. (2013))..... 3
- Figure 1-3** The CO₂ is black, rock is the lightest phase, and brine is intermediate (figure taken from Andrew et al. (2015))..... 4
- Figure 1-4** Geological storage mediums with typical depths (after Thomas and Benson, 2015))..... 5
- Figure 1-5** Conceptualization of in-situ trapping mechanisms (after Burnside & Naylor, 2011). 6
- Figure 1-6** A sketch of key processes governed by capillary trapping after CO₂ injection has ceased at a storage site. Plume migration is limited by the trapping as large fractions of the plume are immobilised. Capillary trapping is secure over long timescales and avoids buoyant stress on overlying cap rock layers. Trapping is also key to parameterising hysteresis in relative permeability functions—more trapping leads to greater disconnection of fluid ganglia as CO₂ saturation in the pore space decreases (movement from A towards C in the figure) and, thus, a larger decrease in permeability as a function of saturation (figure taken from Krevor et al. (2015))..... 8
- Figure 1-7** Forcefield acting on a three-phase mineral/CO₂/brine system. There are three different interfacial forces: solid/CO₂interfacial tension (γ_{sc}), solid/water

interfacial tension (γ_{sw}), and CO₂/brine interfacial tension (γ_{cw}). This figure was taken from Iglauer et al. (2015)..... 9

Figure 1-8 Three-dimensional segmented renderings of HRXCT data showing s-CO₂ saturation after s-CO₂ flood and the trapped s-CO₂ after the brine flood in (a, d) water-wet glass bead pack, (b, e) water-wet angular glass grain pack, and (c, f) CO₂-wet Teflon bead pack. The inset in figure (d) shows individually trapped s-CO₂ in a discrete pore. The scale (see white and black bars) is the same for all images. Solid media are rendered transparent. See the six corresponding animated versions of the images in the supporting information. This figure was taken from Chaudhary et al. (2013)..... 11

Figure 1-9 Layout of thesis objectives and structure 14

Figure 2-1 Pressures acting on a fluid column at the reservoir-seal interface for structural trapping. If CO₂ has migrated upward through the underlying water column, then a fraction of the CO₂ is residually trapped in the pore space of the reservoir rock. Note that at depths below 800 m, CO₂ is in the supercritical state. The insets on the right side schematically illustrate the pore-scale fluid configurations for a water-wet condition. This figure was taken from Iglauer et al. (2015)..... 17

Figure 2-2 Initial-residual saturation relationships from (a) pore-network modelling and (b) laboratory measurements. In the pore network modelling study, the initial residual saturation relationship of oil is shown for a range of contact angles. Most data exhibit water-wet behaviour with the exception of the illite-rich Mt. Simon sandstone and possibly Indiana limestone. This figure was taken from Iglauer et al. (2015)..... 19

Figure 2-3 A schematic of the relative importance of various trapping mechanisms over time (from Benson et al. (2005, 2012)). Residual trapping is significant both in the amount of trapping capacity it provides, as well as for the speed in which residual trapping takes place, simultaneously with water influx into the migrating plume. ... 20

Figure 2-4 CO₂ fate simulation results from the modified tenth SPE Comparative Solution Project model using saturation functions derived from the contact angle to simulation workflow illustrated in Figure 2-5. (a) In the water-wet scenario,

significantly more CO₂ is immobilised as a residually trapped phase than (b) in the CO₂-wet scenario, leading to significantly reduced CO₂ leakage risk. This figure was taken from Iglauer et al. (2015)..... 25

Figure 2-5 An example workflow leading from (a) millimetre-scale contact angle measurements to (d) hectometre-scale pilot-project simulation via (b) millimetre-scale pore-network modelling and (c) the predicted macroscale saturation functions for relative permeability and capillary pressure. The image of a pore network model was reproduced from Blunt et al. (2013). The permeability array is displayed per grid-block on a logarithmic scale in the pilot-scale model (d): the permeability scale is from 0.001 mD (green) to 10,000 mD (red) (9.87×10^{-19} - 9.87×10^{-12} m²). This figure was taken from Iglauer et al. (2015)..... 26

Figure 3-1 Atomic force microscopy images of the dolomite surface used in the experiments. (a) 3D topography of the substrate; (b) deflection signal, different heights are coloured differently (black is 0 nm, white is the maximum). For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article. 32

Figure 3-2 a) advancing and b) receding brine contact angles in the presence of CO₂ measured on hydrophilic dolomite at different pressures (0.1, 5, 10, 15 and 20 MPa) and temperatures (308, 323 and 343K). 38

Figure 3-3 a) advancing and b) receding brine contact angles in the presence of CO₂ measured on hydrophobic dolomite at different pressures (0.1, 5 and 15 MPa) and temperatures (308, 323 and 343K). 39

Figure 4-1 Pressure-Temperature cells for CT: a) pressure cell for 5 mm diameter samples; b) pressure cell for 38.5 mm diameter samples. 1) PEEK pipe; 2) Stainless steel, 3/4" nuts with ferrule; 3) 3/4" to 1/4" unions; 4) confining pressure line 1/8"; 5) pore pressure/ flooding line; 6) 1/4" tees union; 7) stage to be fixed to X-Y-Z stage of the micro-CT. 44

Figure 4-2 (a) Fabrication of the PEEK tube in the lab; (b) cell assembly: sample inside the white rubber sleeve is connected to flooding lines; (c) hydraulic test outside the micro-CT (confining pressure is 33 MPa, zero pore pressure). 45

Figure 4-3 Pressure - Temperature cell inside micro-CY. X-Ray source in on the left, X-Ray objective (detector) is on the right. 45

Figure 4-4 Time lapse of X-ray radiographs of the sample inside the cell at 20 MPa confining pressure: (a) dry sample; (b) brine injection; (c) after CO₂ injection. Low resolution image using 0.4x objective..... 47

Figure 4-5 Tomographic images in the flooding experiment. (a) dry sample; (b) 10%NaI-doped brine saturated sample; (c) after flooding with liquid CO₂. The direction of CO₂ injection is perpendicular to the paper plane. X-ray energy is 80kV, the distance between X-Ray source and the sample is 20 mm, the distance between the sample and 4x objective is 20 mm; pixel size is 3.4348 micrometer. Confining pressure is 20 MPa, pore pressure is 10 MPa, the temperature is 28°C. Note that the images (b) and (c) are corresponding to the same area of the sample. 48

Figure 4-6 Influence of Pressure, Temperature and Fluid on the image quality. Bernheimer sandstone, gas permeability 1880mD. X-Ray energy is 60kV, the distance between X-Ray source and the sample is 19 mm, the distance between the sample and 4x objective is 19 mm; pixel size is 3.4348 micrometer. Number of scans (radiographs) is 4001: (a) Dry sample, Pconf=30 MPa, T=28°C; (b) after liquid CO₂ injection into brine saturated sample, Pconf=25 MPa, Ping=10 MPa, T=28°C; (c) after “near critical” CO₂ injection into brine saturated sample Pconf=15 MPa, Ping=7 MPa, T=42°C; (d) after supercritical-CO₂ injection into brine saturated sample Pconf=15 MPa, Ping=10 MPa, T=42°C. 50

Figure 5-1 μ CT image of Bentheimer sandstone at 10MPa pore pressure and 318K: (a) water-wet initial CO₂ saturation, raw image; (b) water-wet initial CO₂ saturation, segmented image; (c) CO₂ clusters in 3D for the water-wet initial CO₂ saturation, a volume of 3mm³ is shown; (d) water-wet residual CO₂ saturation, raw image; (e) water-wet residual CO₂ saturation, segmented image; (f) CO₂ clusters in 3D for the water-wet residual CO₂ saturation, a volume of 3mm³ is shown; (g) oil-wet initial CO₂ saturation, raw image; (h) oil-wet initial CO₂ saturation, segmented image; (i) CO₂ clusters in 3D for the oil-wet initial CO₂ saturation, a volume of 3mm³ is shown; (j) oil-wet residual CO₂ saturation, raw image; (k) oil-wet residual CO₂ saturation, segmented image; (l) CO₂ clusters in 3D for the oil-wet residual CO₂ saturation, a

volume of 3mm^3 is shown. CO_2 is black/dark grey, brine is grey and sandstone is light grey; in the segmented images CO_2 is yellow, brine is blue and rock is grey... 56

Figure 5-2 Morphologies and volumes of the largest initial and residual CO_2 clusters in Bentheimer sandstone at 318K and 10MPa pore pressure: (a) water-wet initial CO_2 saturation (volume = $27.7 \times 10^6 \mu\text{m}^3$), (b) water-wet residual CO_2 saturation (volume = $5.39 \times 10^6 \mu\text{m}^3$), (c) oil-wet initial CO_2 saturation (volume = $18.4 \times 10^6 \mu\text{m}^3$), (d) oil-wet residual CO_2 saturation (volume = $3.03 \times 10^6 \mu\text{m}^3$)... 58

Figure 5-3 Initial and residual CO_2 droplet size distributions measured in water-wet and oil-wet Bentheimer sandstone at 318K and 10MPa pore pressure..... 59

Figure 5-4 CO_2 droplet surface areas plotted against their volumes for various conditions, at 318K and 10MPa pore pressure..... 60

Figure 5-5 CO_2 (a) and water (b) droplet capillary pressures measured in water-wet and oil-wet Bentheimer sandstone at 318K and 10MPa pore pressure..... 62

Figure 6-1 2D slices through the rock and various fluids at different saturation states. (a) connate water saturation S_{wc} ; (b) residual oil saturation S_{or} ; (c) initial gas saturation $S_{gi,g}$; (d) initial gas saturation $S_{gi,bg}$; (e) residual gas saturation $S_{gr,gb}$; (f) residual gas saturation $S_{gr,bgb}$. Left: greyscale images; right: segmented images. In the raw images, oil is white, gas is black, brine dark grey, and sandstone is light grey. Oil is red, brine blue, gas yellow and rock is brown in the segmented images. 70

Figure 6-2 3D topology of the gas (yellow), oil (red) phases at various saturation states. All volumes shown are $\sim 12\text{mm}^3$ 71

Figure 6-3 Comparison of oil and gas phase configurations at the two residual gas saturation states ($S_{gr,gb}$ and $S_{gr,bgb}$) in water-wet (top) and intermediate-wet (bottom) Clashach sandstone. Note that the plugs were drilled from the same block. The water-wet plug images were acquired with monochromatic light at a nominal resolution of $(9\mu\text{m})^3$ (Iglauer et al. 2013); the intermediate-wet plug images are those discussed in this script, their resolution is $(3.4\mu\text{m})^3$ 74

Figure 6-4 Selected high resolution 2D slices through the plug at residual gas saturation; top: $S_{gr,gb}$, $(1.00\mu\text{m})^3$ resolution; bottom: $S_{gr,bgb}$, $(1.07\mu\text{m})^3$ resolution. Gas

is black, oil white, brine dark grey and rock light grey. For classification consult Table 6-3. 75

Figure 6-5 Surface area versus cluster volume for each oil cluster and production step. 80

Figure 6-6 Surface area versus cluster volume for each gas cluster and production step. 81

Figure 6-7 Oil cluster size distribution for each saturation state..... 82

Figure 6-8 Gas cluster size distribution for each saturation state. 83

Figure 6-9 Oil curvature distribution at various saturation states. 85

Figure 6-10 Gas curvature distribution at various saturation states..... 85

Figure 6-11 Distribution of capillary pressures for oil at various saturation states. 86

Figure 6-12 Distribution of capillary pressures for gas at various saturation states. 87

List of Tables

Table 1-1 Wettability criteria based on the contact angle of rock/CO ₂ /brine systems (modified after Iglauer et al. (2015a)).	10
Table 3-1 Minerals identified in Dolomite sample based on XRD pattern interpretation.	34
Table 5-1 Initial and residual CO ₂ saturations in water-wet and oil-wet Bentheimer sandstone at 318K and 10MPa pore pressure.	56
Table 5-2 Statistical parameters associated with the initial and residual CO ₂ clusters for oil-wet and water-wet sandstone (Bentheimer, 318K, 10MPa pore pressure).	58
Table 6-1 Interfacial tensions of the fluids used.	67
Table 6-2 Porosities, oil S_o , water S_w , and gas saturations S_g measured on the μ CT images (percentages shown). *water-wet condition (taken from Iglauer et al. 2013). R_f is the oil recovery factor, C_{trap} is the capillary trapping capacity.	69
Table 6-3 Pore-corner fluid configurations identified in the μ CT images of the intermediate-wet plug.	76
Table 6-4 Oil and gas cluster statistics: cluster size distribution exponents τ and interfacial area exponents p for intermediate-wet and water-wet sandstone. *Reassessed value: data taken from Iglauer et al. (2014). **Reassessed value: data taken from Iglauer et al. (2013). ***Taken from Iglauer et al. (2013). #4Taken from Pentland et al. (2012).	81
Table 6-5 Best least square fitting equations for the measured oil and gas cluster size distributions in water-wet and intermediate-wet sandstone. *Data taken from Iglauer et al. (2014). **Data taken from Iglauer et al. (2013).	83

Chapter 1 Introduction

1.1 Background

Multiphase fluid flow and displacement in rock porous media is a highly applicable process in the energy and environment fields. Water imbibition processes have been used to displace and produce hydrocarbons, which are trapped by capillary forces in the pore spaces. However, in geological carbon capture and storage (CCS), there are many important trapping mechanisms that are rooted in interfacial phenomena, apparent at the scale of individual pores and throats.

Pore-scale imaging and modelling have been becoming a routine service in the oil and gas industry over the last 10 years because they have potential applications in contaminant transport and carbon dioxide storage; therefore, pore-scale modelling has developed rapidly (Blunt et al. 2013). Furthermore, three-phase fluid flow through porous media is an important topic in the oil and gas industry as it occurs in gas-assisted enhanced oil recovery (EOR) processes and carbon geo-storage.

Petrophysical properties can be used as an indication of the average rock wettability (Anderson, 1986); however, determining the distribution of wettability at the pore scale is complicated. Bakke and Øren (Bakke and Øren, 1997) reconstructed grain-size distribution and other petrographical data from 2D thin sections to 3D sandstones. Note that the reconstructed model reproduced important geometrical and connectivity properties. Valvatne and Blunt (Valvatne and Blunt, 2004) studied the change of wettability at the pore scale using two-phase capillary-dominated network models. They estimated relative permeability and oil recovery for the rock at different wettability. Further, Hui and Blunt (Hui and Blunt, 2000) discussed the effects of rock wettability on three-phase systems after analysing ten different fluid configurations in a single pore and estimating all possible displacements. Hui and Blunt found there is a significant influence of wettability on relative permeability estimation. Moreover, a three-dimensional mixed-wet network model was found by Piri and Blunt (Piri and Blunt, 2005) to simulate two- and three-phase

relative permeability using process-based techniques (Øren and Bakke, 2002) for the pore space in Berea sandstone.

In the oil and gas industry, micro-CT scans have been used to measure pore morphologies for rocks and to visualise porous media, as well as many other static petrophysical parameters (porosity and interface area) and dynamic properties (permeability, NMR response, formation factor, and capillary pressure). Morphologies of carbonate rocks were investigated using X-ray microtomography (Arns et al. 2005). Arns et al. (Arns et al. 2005) identified 32,000 separate vugs and measured the size distribution. In addition, the permeability, drainage capillary pressure, formation factor, and NMR response were estimated using micro-CT images. Cuttings of carbonate core plugs were scanned using micro-CT scans to measure the density and porosity (Siddiqui et al., 2005). Further, an environmental scanning electron microscope (ESEM), NMR spectrometer, and APEX were used for comparison to support pore-scale network modelling. X-ray microtomography was used to get the pore-scale images from rock samples at different resolutions (Blunt et al. 2013; Andrew et al. 2014), see figures 1.1, 1.2, and 1.3 below, which can provide a three-dimensional image of the pore space. However, these images cannot be directly input into network models (Dong and Blunt, 2009). It is, therefore, vital to deepen our understanding in this area to aid in developing better CCS and EOR schemes and to improve clean-up technologies for contaminated soils.

Pore-scale three-phase fluid flow has previously been studied theoretically (e.g. Van Dijke et al. 2011; Piri and Blunt 2004, 2005a, b; Al-Dhahli et al. 2012) or experimentally using 2D micromodels (e.g. Oren 1992, Oren and Pinczewski 1995, Dong et al. 1995, Keller et al. 1997); both approaches rely on strongly simplified pore morphologies. It is, however, well established that pore space morphology has a significant impact on the related fluid dynamics (Dullien 1979). In addition, it is clear that 3D flow characteristics cannot be fully captured by 2D models (e.g. the percolation threshold is significantly lower in 3D, Stauffer 1979). Only recently, with the advent of micro-computed tomography (μ CT), can three-phase flow through porous media be studied in 3D at the pore-scale of realistic geological porous media (compare the recent reviews by Blunt et al. (2013), Wildenschild and Sheppard

(2013), and Cnudde and Boone (2013)), which is most relevant for industrial applications.

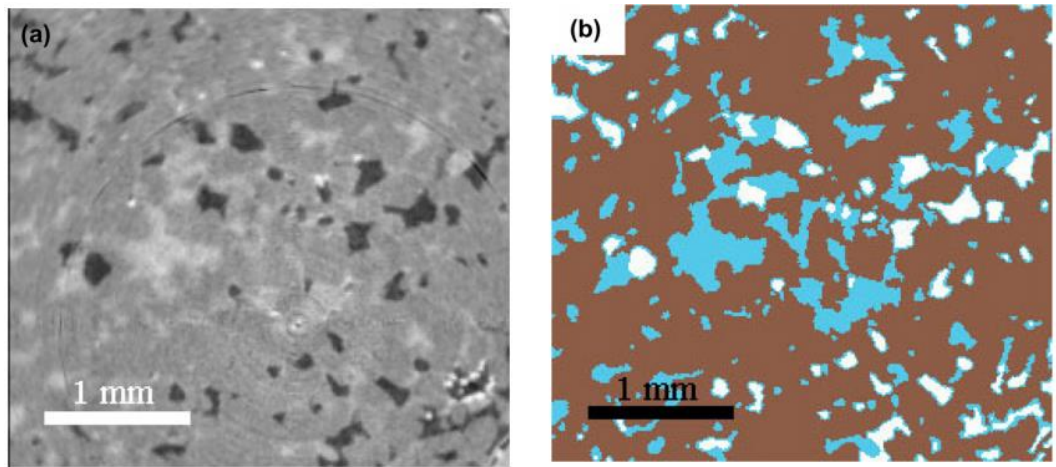


Figure 1-1 Images of CO₂ after waterflooding with CO₂ saturated brine, showing the trapped supercritical CO₂ clusters. The picture is a two-dimensional cross-section through a three-dimensional image; the area shown is 10.80 mm² (261 239 voxels). (a) A raw grey-scale image: supercritical CO₂ is black, brine is light grey, and sandstone is dark grey. The few white areas are minerals with high X-ray absorption. (b) The same slice with phases segmented. Supercritical CO₂ is white, brine light blue, and rock is brown (figure taken from Blunt et al. (2013)).

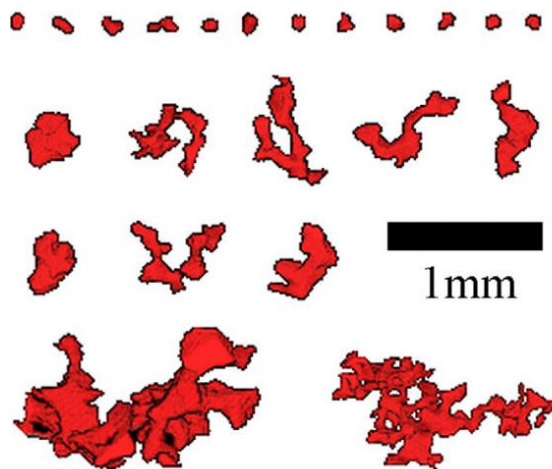


Figure 1-2 Images of CO₂ after waterflooding with CO₂ saturated brine showing selected trapped clusters of different sizes (figure taken from Blunt et al. (2013)).

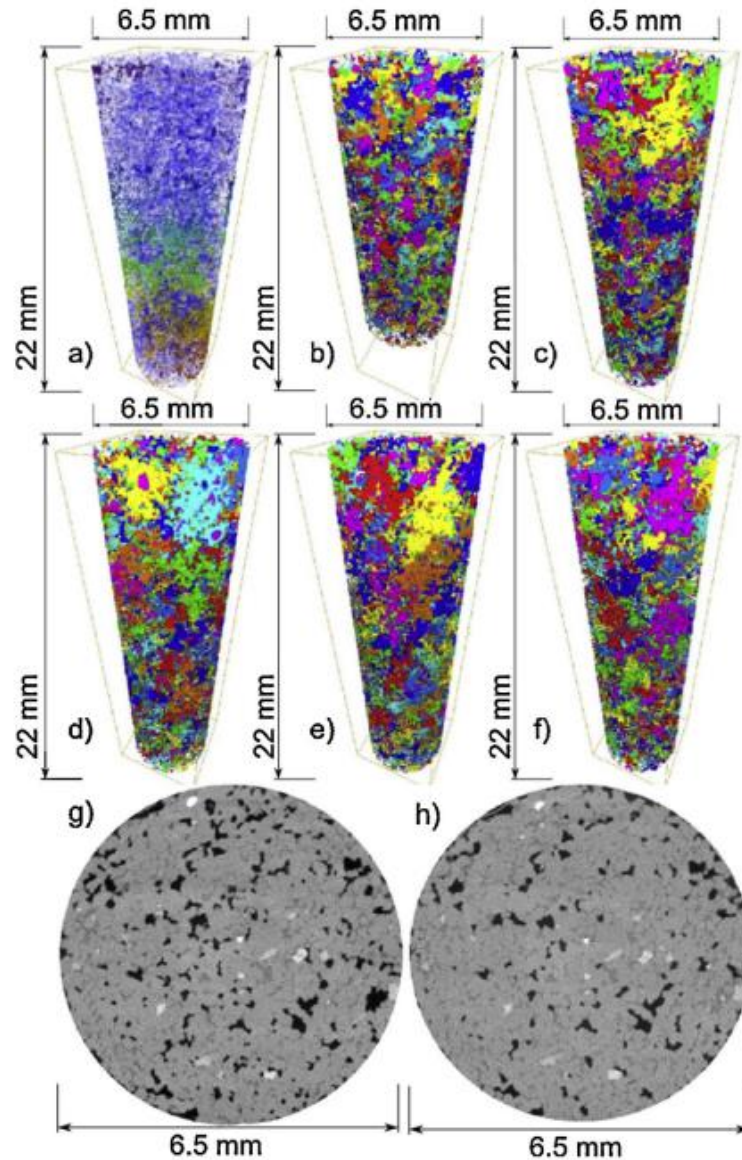


Figure 1-3 The CO₂ is black, rock is the lightest phase, and brine is intermediate (figure taken from Andrew et al. (2015)).

In the last few decades, high concentration of carbon dioxide emissions in our atmosphere has become a major issue worldwide as it leads to global warming and climate change. CO₂ emissions have dramatically increased since the industrial revolution (Iglauer et al. 2012) owing to increasing demand for energy. Fossil fuels are the major source of energy (International Energy Agency, 2010). They contribute to >80% of total energy consumption. However, fossil fuels are depleting at an increasing rate and this situation challenges us to develop advanced technologies to recover additional fossil fuels.

Further, anthropogenic CO₂ emissions can be cut by injecting CO₂ deep underground. This method, known as carbon capture and storage (CCS) (Metz et al. 2005), has the potential to significantly reduce CO₂ emissions in the atmosphere. Crude oil is the most important fossil fuel and is only recovered up to ~30% (BP Statistical Review of World Energy, 2014). However, several methods can be used to enhance recovery (EOR), such as miscible or partially miscible gas flooding, thermal stimulation (Lake, 2010), polymer flooding (Iglauer et al. 2010), and surfactant flooding (Iglauer et al. 2010). CO₂ enhanced oil recovery is attractive (Firoozabadi et al. 2010; Blunt et al. 1993), as it can be combined with CO₂ sequestration.

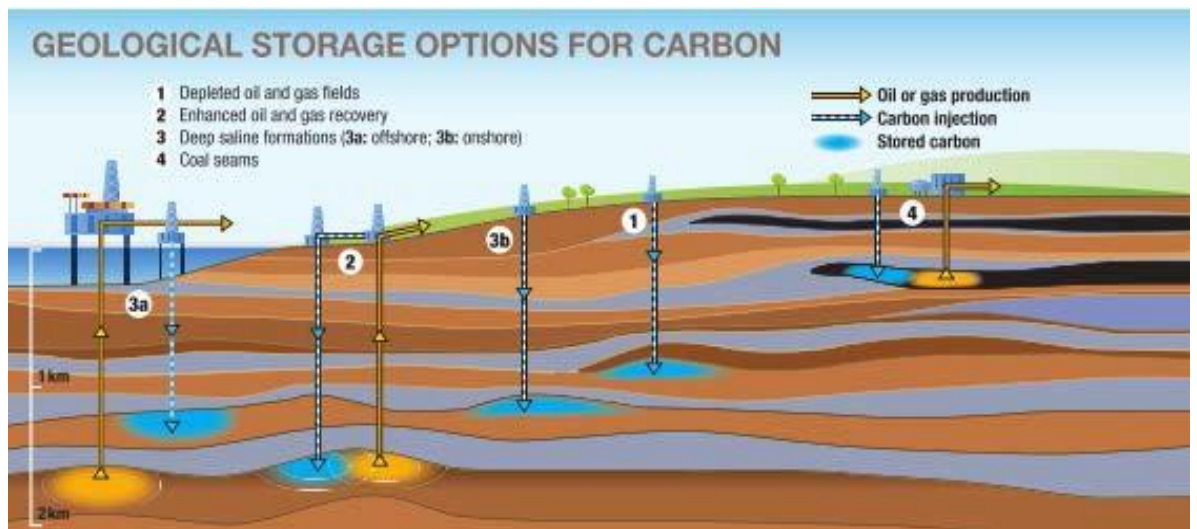


Figure 1-4 Geological storage mediums with typical depths (after Thomas and Benson, 2015)).

Residual trapping (Figure 1-5) refers to the physical immobilization of CO₂ in the pore spaces because of capillary forces that hold the buoyant CO₂ back (Hesse et al. 2008; Juanes et al. 2006; Pentland et al. 2011). Capillary trapping, which is a rapid and effective mechanism, occurs at the cessation of CO₂ injection into the targeted formation while the brine displaces and traps CO₂ as discontinuous droplets at the trailing edge of a rising CO₂ plume. Further, it was found that rock wettability (Farophpoor et al. 2013; Iglauer et al. 2013; Iglauer et al. 2015a) and interfacial tension

(Chalbaud et al. 2009) significantly influence capillary trapping. It was also reported that initial gas saturation, pore geometry, pore throat aspect ratio, and hysteresis can affect capillary trapping (Al-Menhali et al. 2015; Juanes et al. 2006; Suekane et al. 2010), Figure 1-5.

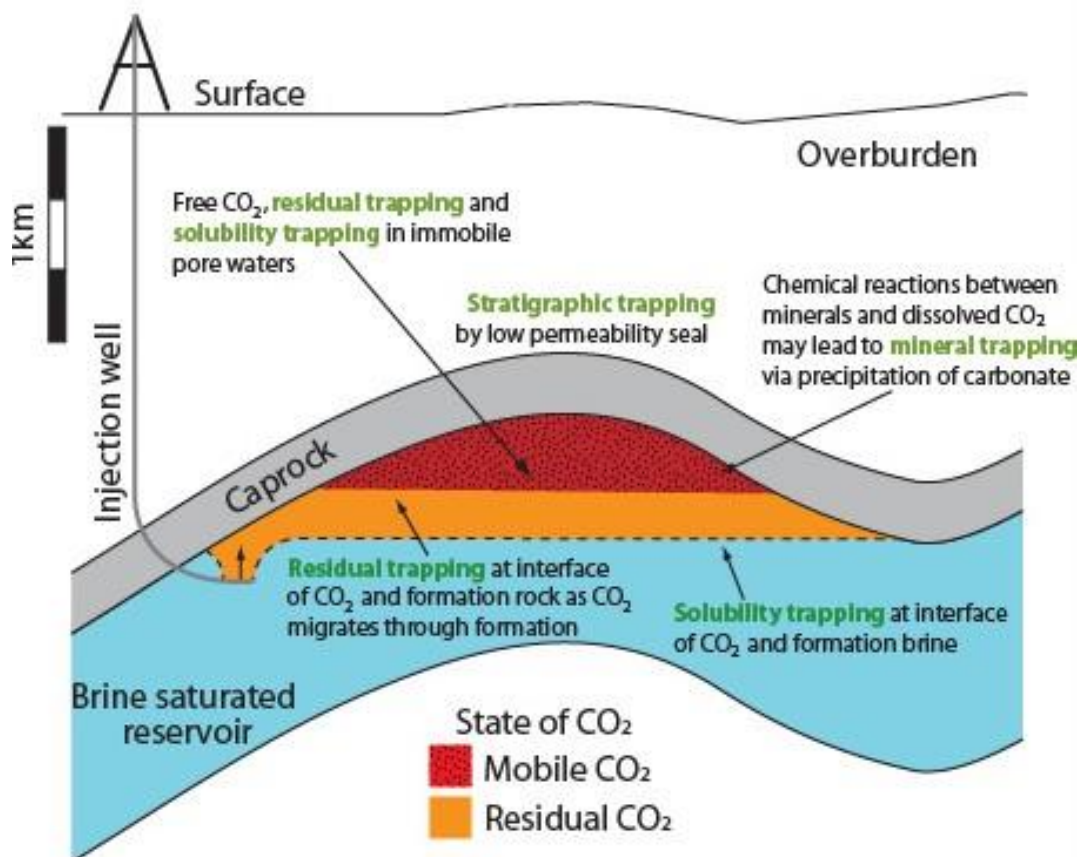


Figure 1-5 Conceptualization of in-situ trapping mechanisms (after Burnside & Naylor, 2011).

However, it is important to understand the nature of interaction between fluid and rock in order to safely store CO₂ in the ground and assess the CO₂'s ability to increase the recovery (Oomole et al. 1983; Iglauer et al. 2015, 2015a). Therefore, the wettability of a rock–CO₂–water system was measured here because it is a key parameter in geological carbon capture and storage (CCS) owing to its impact on the three most important trapping mechanisms during the initial storage phase, i.e.

structural (Iglauer et al. 2015; Naylor et al. 2011), residual (Iglauer et al. 2011, 2016; Krevor et al. 2015; Chaudhary et al. 2013; Rahman et al. 2016; Al-Khdheawi et al. 2016, 2017, 2017a), and dissolution trapping (Al-Khdheawi et al. 2016, 2017, 2017a) (see Figure 1-6).

Anderson (1986 and 1987) and Salathiel (1973) described reservoir rock wettability measurement methods comprehensively by categorising the measurements into direct and indirect methods. Contact angle measurements represent the direct method, which refers to the relative amount of spontaneous (imbibition) and forced displacement (Amott, 1959). Relative permeability (Owens and Archer, 1970), capillary pressure (Slobod and Blum, 1952), flotation, glass slide, and NMR methods represent the indirect method. However, no satisfactory method exists for the in-situ measurement for the rock wettability. Therefore, it is important that it is measured in the laboratory.

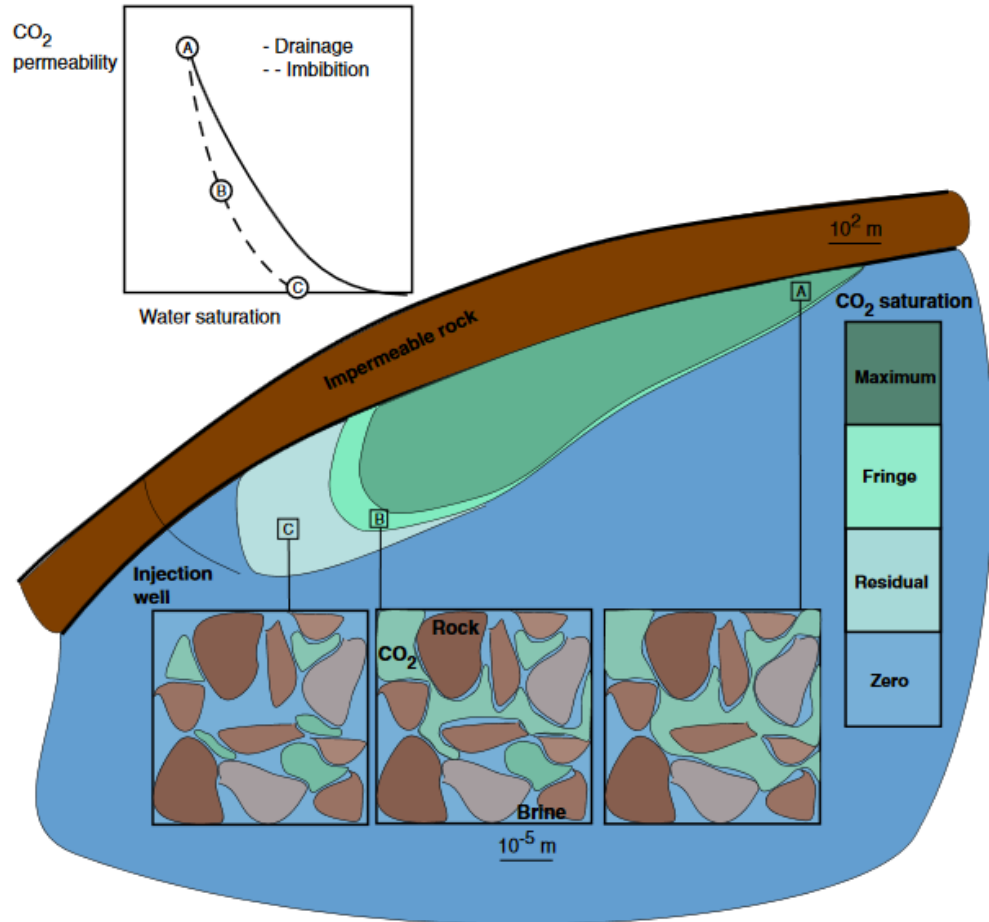


Figure 1-6 A sketch of key processes governed by capillary trapping after CO₂ injection has ceased at a storage site. Plume migration is limited by the trapping as large fractions of the plume are immobilised. Capillary trapping is secure over long timescales and avoids buoyant stress on overlying cap rock layers. Trapping is also key to parameterising hysteresis in relative permeability functions—more trapping leads to greater disconnection of fluid ganglia as CO₂ saturation in the pore space decreases (movement from A towards C in the figure) and, thus, a larger decrease in permeability as a function of saturation (figure taken from Krevor et al. (2015)).

The contact angle measurement (Figure 1-7) is the best quantitative method for characterising wettability, especially when pure mineral samples are used (Giraldo et al. 2013; Mittal, 2006; Kaveh et al. 2012; 2014). The captive bubble method and sessile drop method (Anderson, 1986) are the most popular in the petroleum industry. In that study, the tilting plate method—part of the sessile drop method—

was used because advancing and receding contact angles could be measured simultaneously (Lander et al. 1993). Practically, for a rock/CO₂/brine system, if θ is less than 50°, then the system is strongly water-wet, weakly water-wet for θ between 50-70°, and intermediate-wet for the θ range 70-110°. The system turns weakly CO₂-wet for the θ range 110-130° and strongly CO₂-wet for the θ range 130-180° (Table 1-1). However, pressure, temperature, brine salinity surface roughness, and many other parameters can influence wettability measurements.

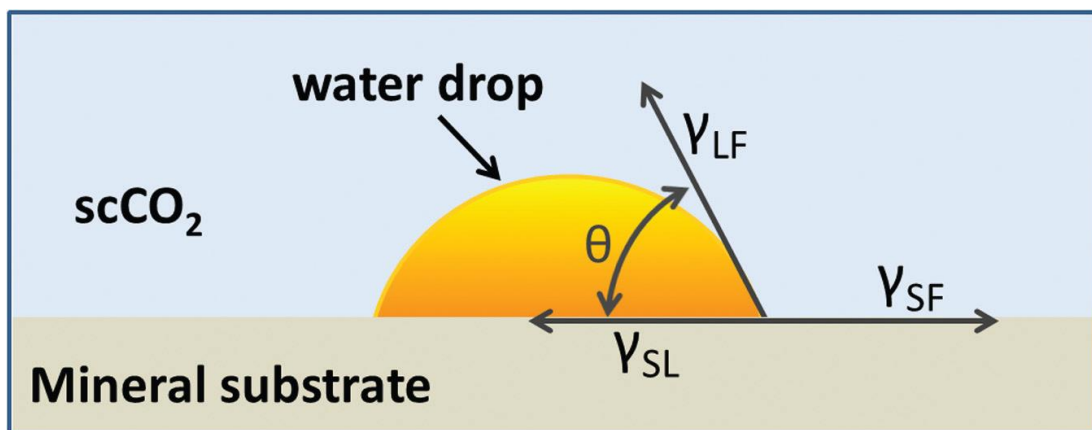


Figure 1-7 Forcefield acting on a three-phase mineral/CO₂/brine system. There are three different interfacial forces: solid/CO₂ interfacial tension (γ_{se}), solid/water interfacial tension (γ_{sw}), and CO₂/brine interfacial tension (γ_{cw}). This figure was taken from Iglauer et al. (2015).

Table 1-1 Wettability criteria based on the contact angle of rock/CO₂/brine systems (modified after Iglauer et al. (2015a)).

Wettability state	Water contact angle θ (°)
Complete wetting	0
Strongly water-wet	0 – 50
Weakly water-wet	50 – 70
Intermediate-wet	70 – 110
Weakly CO ₂ -wet	110 – 130
Strongly CO ₂ -wet	130 – 180
Complete non-wetting of water	180

1.1.1 Influence of wettability on residual trapping

Capillary trapping, which is proportional to the product of porosity times the residual saturation occurred when the capillary forces inside the porous medium trapped the non-wetting phase clusters. Principally, capillary trapping occurs during the imbibition process, i.e. at the trailing edge of the CO₂ plume. This residual CO₂ saturation, which is a function of wettability, is resolved through laboratory measurements of relative permeability and capillary pressure curves (Krevor et al. 2012; Pini and Benson, 2013) or through pore-scale models. Reservoir-scale simulation and 3D micro-CT imaging detect that water-wet rocks have a better residual trapping potential, which refers to trapping large residual/nonwetting phase clusters trapping by snap-off phenomenon, which occurs only in water-wet rocks (Chaudhary et al. 2013; Iglauer et al. 2013; Krevor et al. 2012; Krevor et al. 2015) (Figure 1-8).

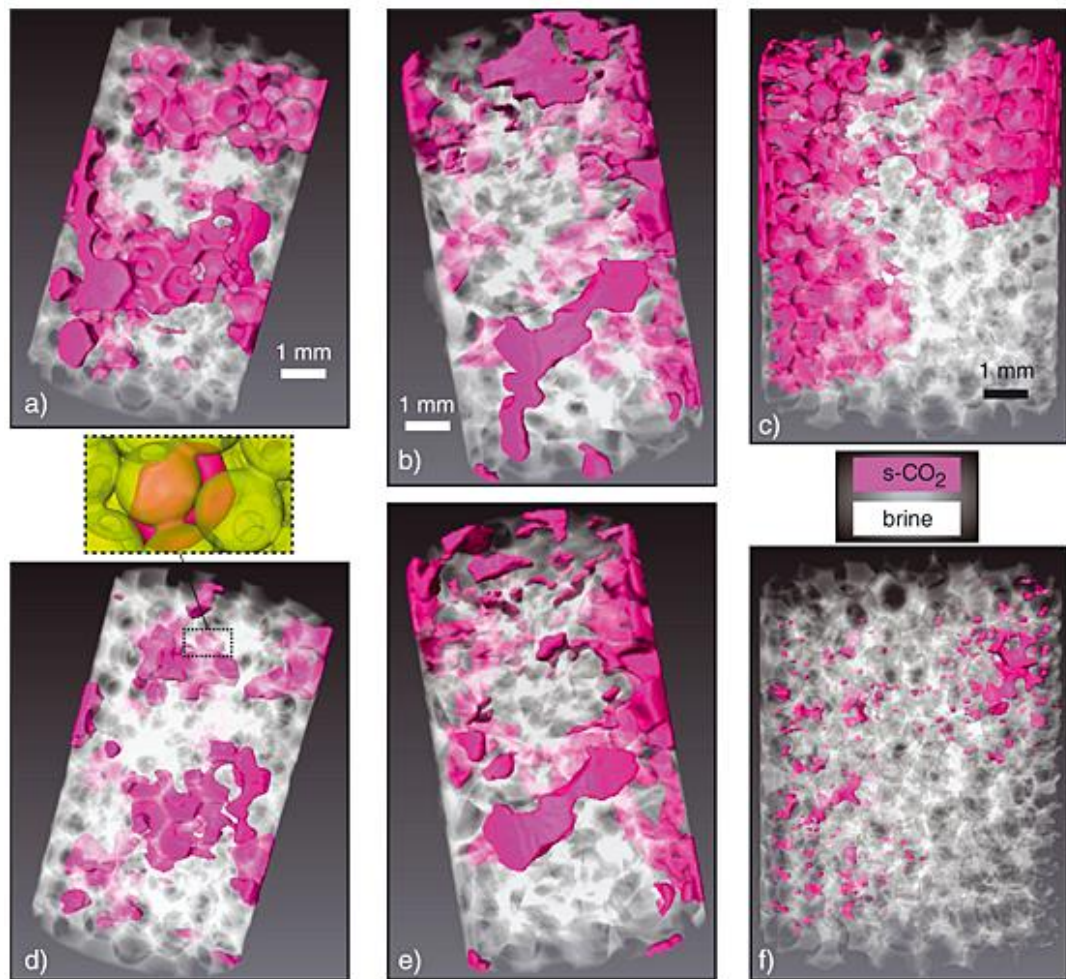


Figure 1-8 Three-dimensional segmented renderings of HRXCT data showing s-CO₂ saturation after s-CO₂ flood and the trapped s-CO₂ after the brine flood in (a, d) water-wet glass bead pack, (b, e) water-wet angular glass grain pack, and (c, f) CO₂-wet Teflon bead pack. The inset in figure (d) shows individually trapped s-CO₂ in a discrete pore. The scale (see white and black bars) is the same for all images. Solid media are rendered transparent. See the six corresponding animated versions of the images in the supporting information. This figure was taken from Chaudhary et al. (2013).

1.2 Research objective

The main objective of this research is to investigate pore-scale displacement processes associated with CO₂-EOR in rocks. This specifically includes the analysis of rocks' microstructure, which is a key factor in determining these flow processes. Various rock samples with different wettabilities were used. We analyse the capability of CO₂ trapping in oil reservoirs that is combined with enhanced oil recovery. We, thus, imaged residual CO₂ supercritical in oil-wet sandstone plug using high-resolution X-ray microcomputed tomography at reservoir conditions.

Specifically, the objectives of this research are:

1. Observing and describing pore-scale oil-CO₂-water displacement processes in rock.
2. Investigating the wettability of rocks influencing oil production and CO₂ sequestration.
3. Analysing the capability of rocks in trapping the residual of CO₂, using X-ray microcomputed tomography.

1.3 Thesis organisation

There are seven chapters in this thesis, starting with the introduction, literature review, results, and conclusion.

Chapter 1. Introduction

This first chapter presents the introduction of the background, the importance of understanding multiphase flow, describes pore-scale oil-CO₂-water displacement processes, and issues related to CO₂ sequestration and the solution. The objectives and thesis organisation are also included in this chapter.

Chapter 2 – Literature review.

In this second chapter, a brief summary of pore-scale analysis and the influence of wettability in residual CO₂ trapping processes is provided.

Chapter 3 – Effect of Temperature on CO₂/Brine/Dolomite Wettability: Hydrophilic vs Hydrophobic Surfaces

This chapter analyses the effect of temperature on CO₂/brine/dolomite wettability using two different types of surfaces.

Chapter 4 – Residual Trapping of Supercritical CO₂: Direct Pore-Scale Observation Using a Low-cost Pressure Cell for Microcomputer Tomography

This chapter describes the design and fabrication steps of a low-cost, modular experimental cell for flooding and X-ray tomography visualisation of the rock samples at reservoir conditions.

Chapter 5 – Residual Trapping of Supercritical CO₂ in Oil-Wet Sandstone

This chapter unravels the capability of the CO₂ residual trapping of CO₂ on water-wet rocks. CO₂ cluster morphologies and sizes are also presented in this chapter.

Chapter 6 – Influence of Wettability on Residual Gas Trapping and Enhanced Oil Recovery in Three-Phase Flow: A Pore-Scale Analysis Using Microcomputed Tomography

This chapter involves observing the capillary pressure, curvature, and residual saturation of CO₂ and oil phases in homogenous sandstone rocks.

Chapter 7 – This last chapter underlines the meaningfulness of the findings in this research study.

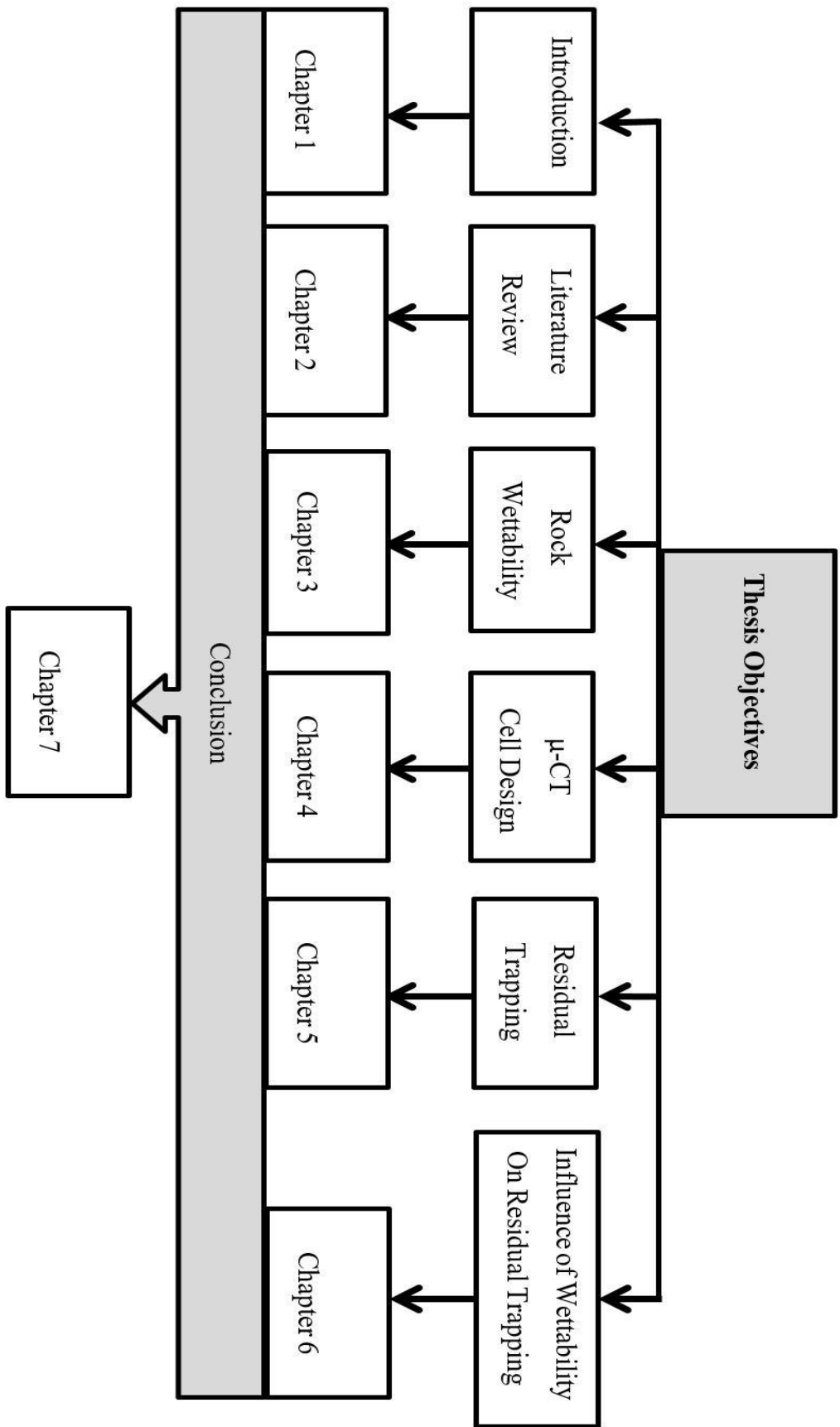


Figure 1-9 Layout of thesis objectives and structure

Chapter 2 Literature Review

2.1 Residual Trapping of Supercritical CO₂

CO₂ injection into the pore network of the hydrocarbon reservoir is a typical process to enhance oil recovery (Olenick et al. 1993; Babadagli 2006) and CO₂ storage (Cantucci et al. 2009), which has been recently recognised as a major cause of climate change. CO₂ can be injected into deep saline aquifers for carbon geo-sequestration (CGS) only, which is a viable option for reducing CO₂ emissions (IPCC 2005, Lackner 2003). Further, it has been reported that deep saline sandstone formation is mostly water-wet even after CO₂ injection (Kaveh et al. 2014; Sarmadivaleh et al. 2015; Al-Yaseri et al. 2016) while most of the hydrocarbon reservoir rock is oil-wet or intermediate-wet (Hirasaki et al. 2004; Cantucci et al. 2009; Arif et al. 2016). However, most existing oil reservoirs have legal (permission from the related authorities to use the site) and supporting infrastructures already in place (Gurevich et al. 2014). Therefore, oil reservoirs have attracted considerable attention in recent years for CO₂ injection (Blunt 1993, Orr 2009, Lake 2010, and Iglauer 2012). On the other hand, pore-scale residual trapping of CO₂ in oil-wet reservoirs has not yet been proven (Cantucci et al. 2009). Moreover, safely trapping CO₂ is greatly required in the sequestration process because it can prevent the CO₂ from leaking to the surface. In addition, the trapping capability, which refers to how much CO₂ can be securely stored as a residual phase without reliance on cap rock seals is important in CCS projects (Iglauer et al. 2009; Iglauer et al. 2011b; Iglauer et al. 2015a).

In saline sandstone formations, which are mostly water-wet, CO₂ would be trapped when it is displaced by the aquifer's brine water (Ennis-King and Peterson, 2005; Kumar et al. 2005). However, in oil reservoirs, which are mostly oil-wet, the effective movement of CO₂ is lower. However, this allows for more of the formations to be contacted by CO₂ while the chase brine rapidly and effectively traps the CO₂ (Qi et al. 2009). Moreover, Al-Menhali et al. (2016) used cylindrical small samples (5 mm diameter and 15 mm length) of Estailades limestone (97.9% calcite and 2.1 quartz) with average permeability of 140 mD and 25% porosity. The

measurements were tested on strongly water-wet and intermediate-wet (Arabian medium crude oil was used for wettability alteration) samples. CO₂+brine or N₂+brine (7 wt% KI brine) was injected at 10 MPa pore pressure, 13 MPa confining pressure, and a temperature of 298 K for N₂ and 323 K for CO₂. Then, the micro CT (μ CT) scanner (Versa XRM-500 X-ray microscope, Zeiss X-ray Microscopy, CA, USA) was used for imaging at (4.9 μ m)³ (resolution). They reported an average initial CO₂ saturation of 29% pore volume, which was similar for both water-wet and intermediate-wet samples. However, the remaining CO₂ after water flooding in the intermediate-wet sample (14%) was significantly lower than the saturation in the water-wet sample (21%). Moreover, the initial and residual N₂ saturations in the water-wet sample were 33% and 24%, respectively, and 20% and 13% for the intermediate-wet sample. Further, a homogenous sandstone (80 mD permeability and 11.1% porosity) outcrop (96 wt% quartz with a small amount of k-feldspar, ankerite and calcite) was selected by Iglauer et al. (2013) for their experiments, which was assumed to be water-wet rock. N₂ gas, brine (10 wt% KI), and oil (bromododecane) were used as fluid phases to study the efficiency of oil recovery using μ CT. Two GEOR flooding sequences were also used for comparison: 1) gas was directly injected into a core at connate water saturation followed by a chase brine injection (gw sequence); 2) gas was injected into a water flooded core at residual oil saturation followed by chase brine injection (wggw sequence). The μ CT scanner was used for imaging at (9 μ m)³. Their results showed that more oil can be produced in the gw process ($S_{or} = 21.6\%$) compared to the wggw process ($S_{or} = 29.3\%$). In 2011, Iglauer et al. also ran their experiment in a homogeneous Doddington sandstone (98 wt% quartz, 2 wt % K-feldspar and traces of kaolinite) sample (4.95 mm diameter and 9 mm length) with a porosity of 20.7% and a brine permeability of 1.6 Darcy. The core sample was initially fully saturated with brine, then CO₂ was injected and finally followed by live brine (10 wt% KI brine saturated with CO₂) injection at 9 MPa pore pressure, 11.75 MPa confining pressure, and temperature of 323 K. The core was then scanned with μ CT at nominal resolution of (13.156 μ m)³, which showed a low CO₂ residual saturation (24.9 % pore volume) for a strongly water-wet system.

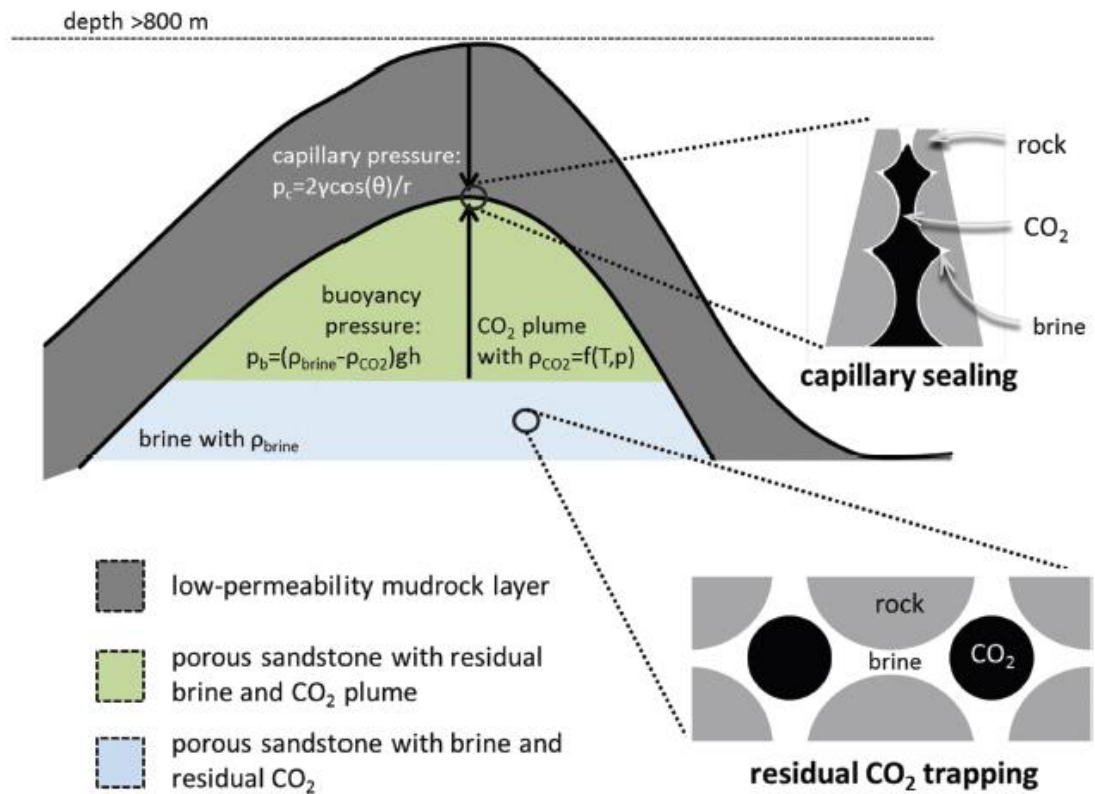
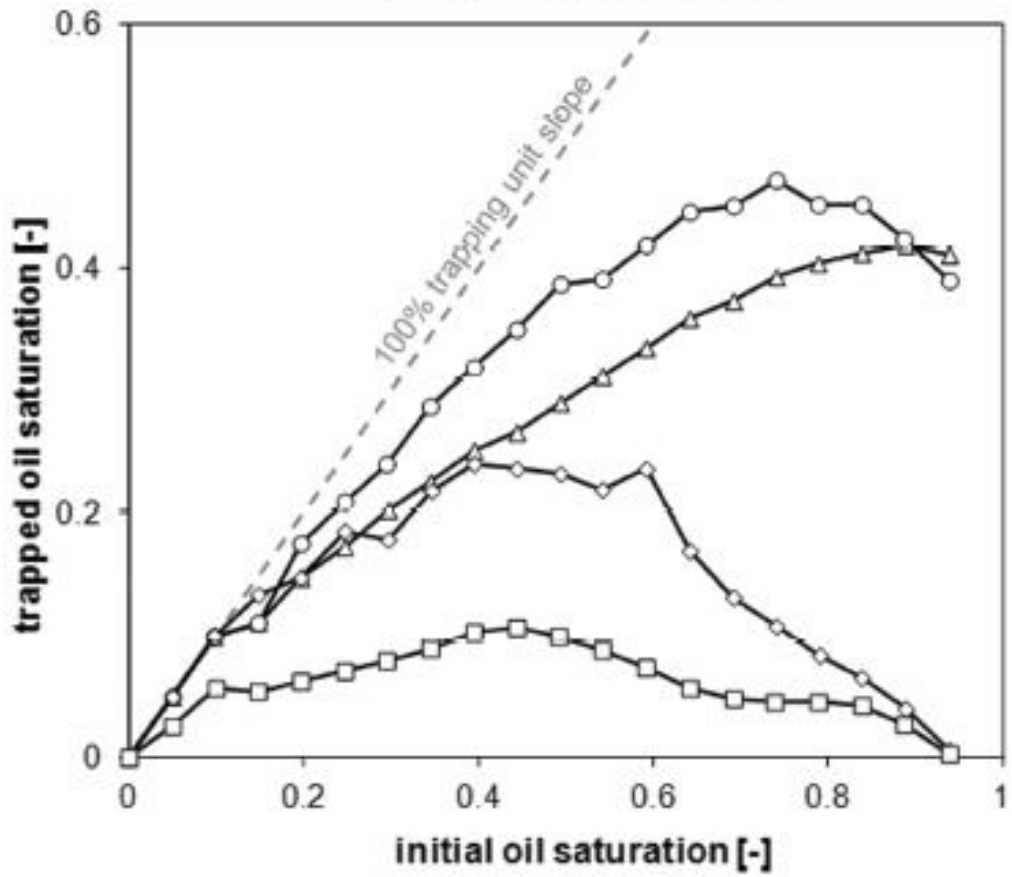


Figure 2-1 Pressures acting on a fluid column at the reservoir-seal interface for structural trapping. If CO₂ has migrated upward through the underlying water column, then a fraction of the CO₂ is residually trapped in the pore space of the reservoir rock. Note that at depths below 800 m, CO₂ is in the supercritical state. The insets on the right side schematically illustrate the pore-scale fluid configurations for a water-wet condition. This figure was taken from Iglauer et al. (2015).

a

- △- $\theta = 20^\circ$ - Spiteri et al., 2008
- $\theta = 60^\circ$ - Spiteri et al., 2008
- ◇- $\theta = 100^\circ$ - Spiteri et al., 2008
- $\theta = 140^\circ$ - Spiteri et al., 2008



b

- ▲ Berea sandstone - Pentland et al., 2011
- ▲ Berea sandstone - Krevor et al., 2012
- ▲ Mt. Simon sandstone - Krevor et al., 2012
- ▲ Paaratte sandstone - Krevor et al., 2012
- ▲ Tuscaloosa sandstone - Krevor et al., 2012
- ▲ Berea sandstone - Akbarabadi and Piri, 2013
- ▲ Nuggett sandstone - Akbarabadi and Piri, 2013
- Indiana limestone - El-Maghraby and Blunt, 2013

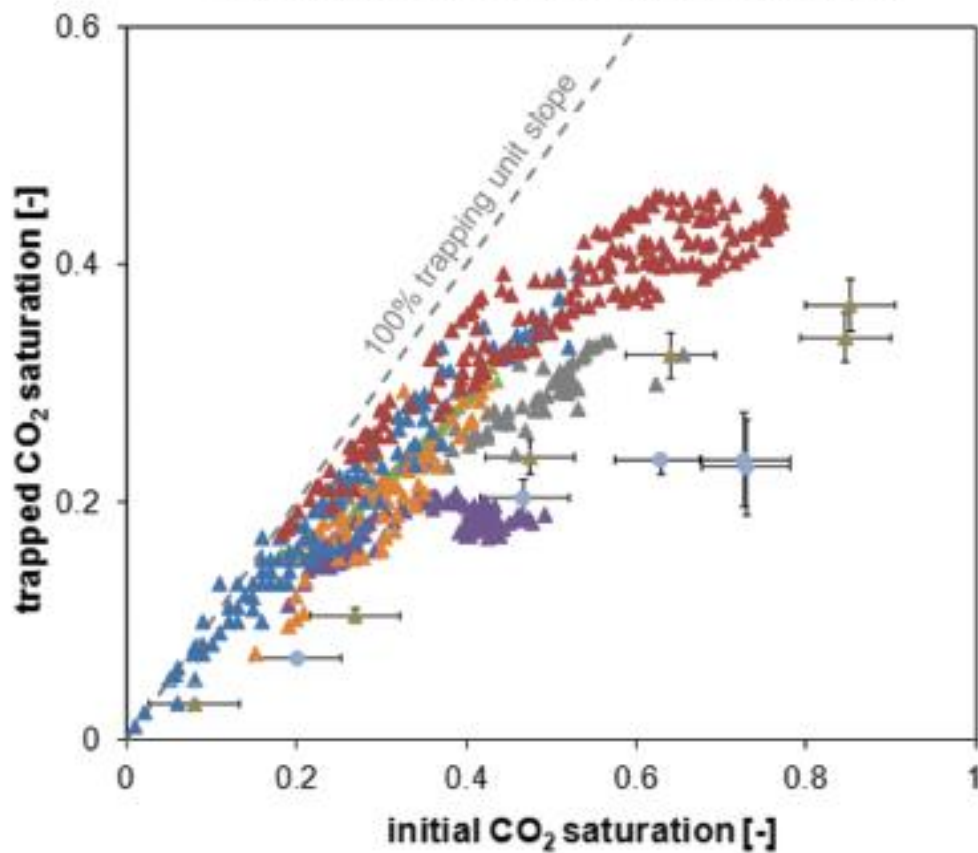


Figure 2-2 Initial-residual saturation relationships from (a) pore-network modelling and (b) laboratory measurements. In the pore network modelling study, the initial residual saturation relationship of oil is shown for a range of contact angles. Most data exhibit water-wet behaviour with the exception of the illite-rich Mt. Simon sandstone and possibly Indiana limestone. This figure was taken from Iglauer et al. (2015).

2.1.1 Storage Mechanisms

There are various physical and chemical mechanisms for storing CO₂ in the porous medium. These include structural trapping, residual trapping, solubility trapping, mineral trapping, and adsorption trapping (Figure 2-3).

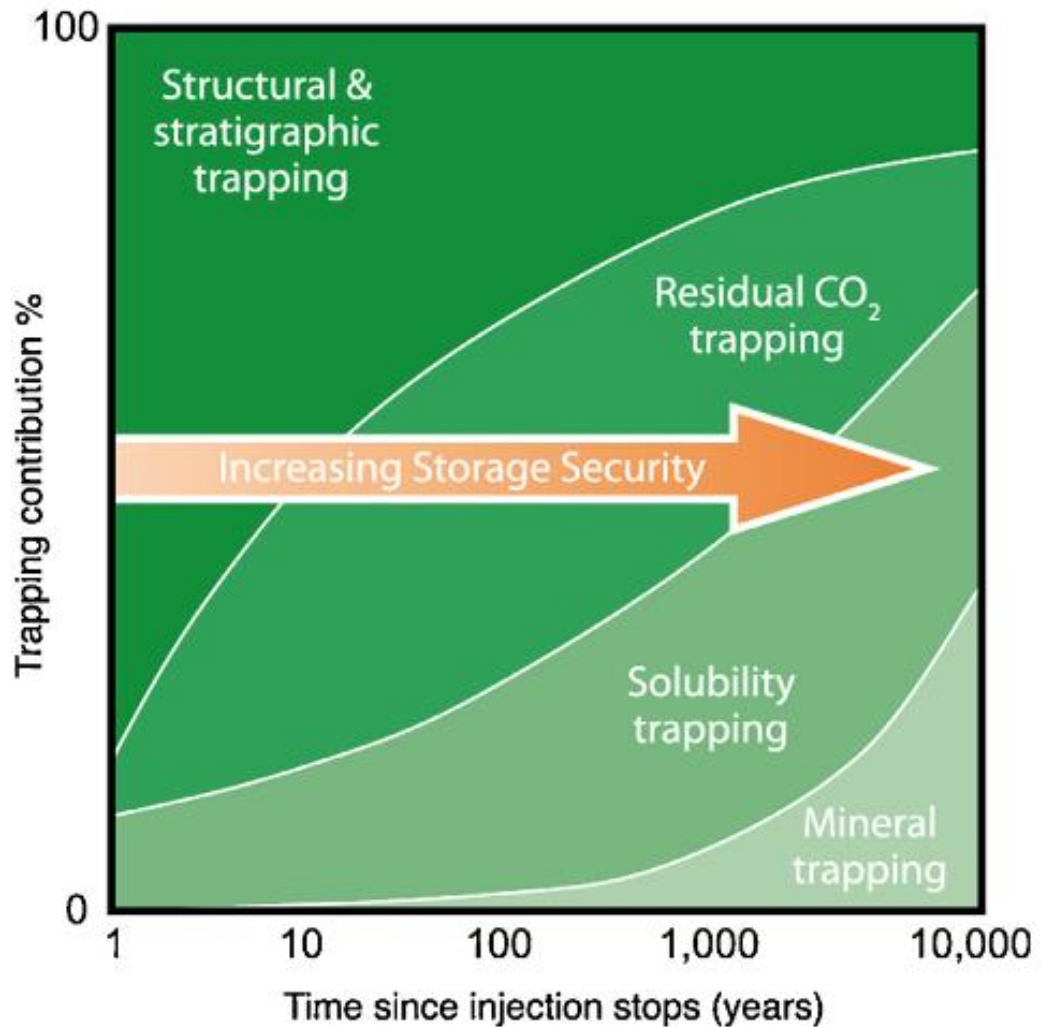


Figure 2-3 A schematic of the relative importance of various trapping mechanisms over time (from Benson et al. (2005, 2012)). Residual trapping is significant both in the amount of trapping capacity it provides, as well as for the speed in which residual trapping takes place, simultaneously with water influx into the migrating plume.

2.1.1.1 Structural trapping

CO₂ will rise after it is injected into porous formations because of the imbalance of buoyancy and capillary forces. Where buoyancy forces exceed capillary forces, the CO₂ will be retained into the porous medium. Meanwhile, the CO₂ plume migrates up-structure and becomes thick and fine-textured in low permeability rock (cap rock) (Neuzil, 1994). This low permeability cap rock or seal prevents CO₂ from upward migration and then CO₂ is structurally trapped. Therefore, structural CO₂ trapping is one of the most dominant storage mechanisms because it can contribute to the total CO₂ trapped volume throughout the storage project life and is most significant in the first decade. Note, it was found that if injection pressure of CO₂ exceeds the fracture pressure of the cap rock, cap rock would fail, which would lead to reduced cap rock sealing efficiency (Chiquet et al. 2007) and the effective storage volume would be limited (Underschultz and Strand, 2016). Further, the rock wettability and interfacial tension can control the capillary sealing efficiency of cap rock (Chiquet et al. 2007; Broseta et al. 2012; Iglauer et al. 2015a, b; Yang et al. 2007).

2.1.1.2 Residual trapping

Residual trapping is the physical immobilisation of CO₂ in the pore spaces because of capillary forces that hold the buoyant CO₂ back (Hesse et al. 2008; Juanes et al. 2006; Pentland et al., 2011). Capillary trapping is one of the most effective storage mechanisms that occurs at the cessation of CO₂ injection to storage formation while the brine displaces and traps CO₂ as discontinuous droplets at the trailing edge of a rising CO₂ plume. The residual CO₂ estimation represents the capillary trapping potential (Kumar et al. 2005; Qi et al., 2009; Saadatpoor et al. 2010) in laboratory practice because high pressure is used to inject brine through the core held. Moreover, capillary trapping is an effective mechanism for a CO₂ storage project. Note that capillary trapping can be demonstrated by wettability (Farophpoor et al. 2013; Iglauer et al. 2013; Iglauer et al. 2015a) and interfacial tension (Chalbaud et al. 2009). Capillary trapping is also influenced by pore throat aspect ratio, pore geometry, initial gas saturation, and hysteresis (Al-Menhali et al. 2015; Juanes et al. 2006; Suekane et al. 2010).

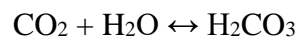
2.1.1.3 Solubility trapping

The dissolution of CO₂ in brine formation leads to a higher density of CO₂ saturated brine, which sinks to the bottom of the reservoir and can be referred to as solubility trapping. Solubility trapping is a geochemical storage mechanism that has several benefits. Dissolved CO₂ does not exist as a separate phase. Thereby, upwards CO₂ plume migration owing to buoyancy is non-existent (Iglauer, 2011c; IPCC, 2005). However, solubility trapping is controlled by operating pressure, temperature, and brine salinity (Koschel et al. 2006; Lagneau et al. 2005) specifically from mid to long-term scales (Gilfillan et al. 2009).

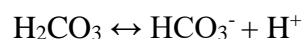
2.1.1.4 Mineral trapping

Mineral trapping or mineralization is an important mechanism for long timescales (thousands of years), where carbon dioxide is dissolved in formation brine, resulting in a weak carbonic acid (Bachu et al. 1994; Xu et al. 2001) that can react with different minerals inside the reservoir rock (carbonate or silicate). However, this will not affect storage security during the CO₂ injection project because the time scale of this trapping needs thousands of years. The following reaction equations explain the mineral- trapping process (Bachu et al. 1994; Xu et al. 2001):

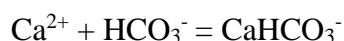
First, the injected CO₂ dissolves in formation brine forming carbonic acid as depicted by the following chemical reaction:



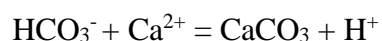
Next, the carbonic acid undergoes a rapid dissociation to form the bicarbonate ions as follows:



Then the increased acidity promotes dissolution of the primary minerals in the host rock, which results in a bicarbonate complex as follows:



Finally, the dissolved bicarbonates react with the divalent mineral cations to result in carbonate precipitation, e.g. the formation of calcium carbonate as follows:

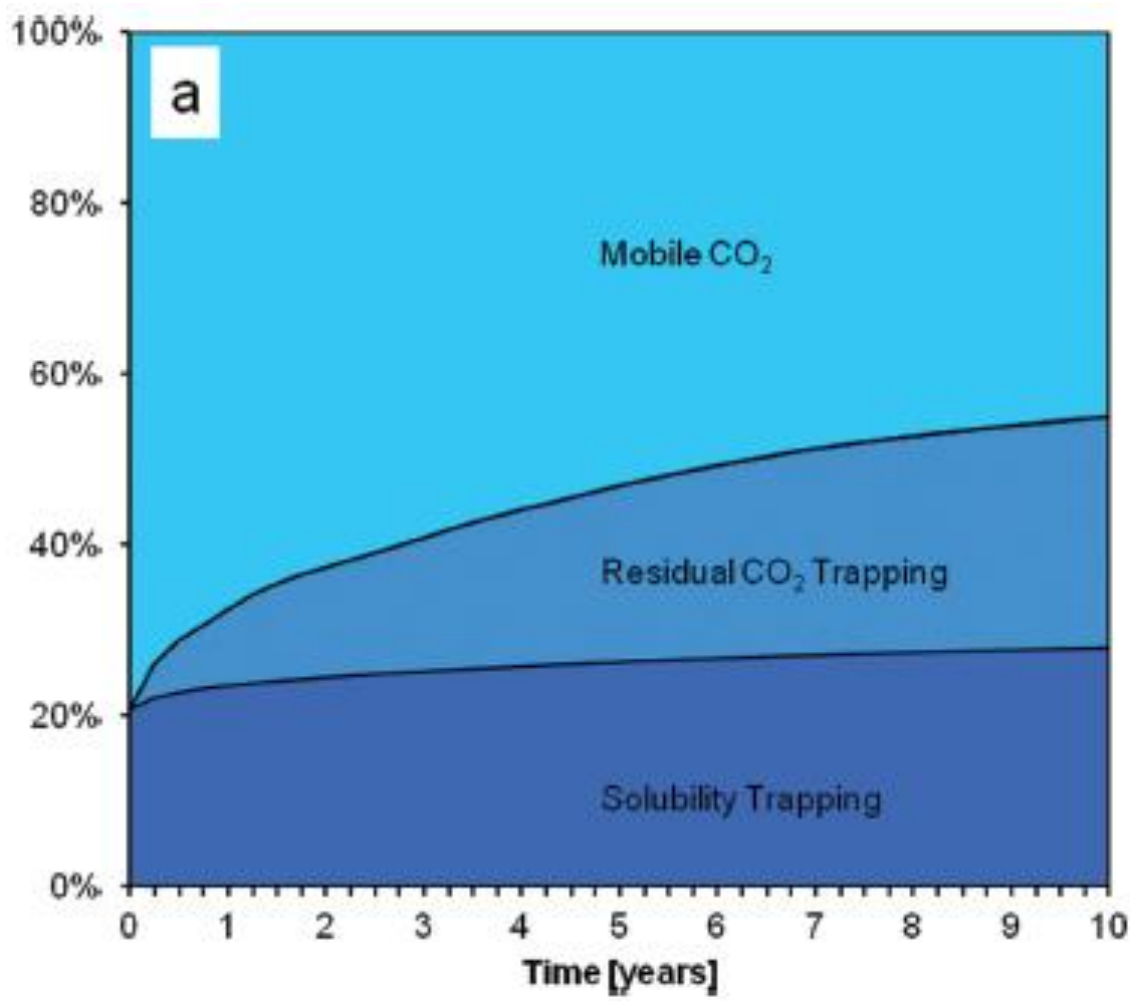


2.1.1.5 Adsorption trapping

This type of mechanism is only applicable in coal bed CO₂ storage by which means the injected CO₂ will displace attached methane from coal because CO₂ has a greater affinity to coal than methane (Busch et al., 2004; Clarkson and Bustin, 2000). However, coal has a very small storage capacity.

2.2 Influence of Wettability on Residual Gas Trapping and EOR

Three-phase flow through porous media is an important topic in the oil and gas industry as it occurs in gas-assisted enhanced oil recovery (EOR) processes, and carbon geo-storage (CCS), when CO₂ is injected into oil reservoirs for storage and incremental oil recovery (Blunt et al. 1993, IPCC 2005). Further, three-phase flow also occurs in soils contaminated by organic solvents (Sleep and McClure 2001). In this context, a significant amount of work has been completed to study such three-phase flow at the reservoir scale (e.g. compare Qi et al. 2009; Moortgat et al. 2011; Beygi et al. 2013) and the mesoscale (centimetres up to several metres sized structures) (e.g. Oak et al. 1990; Egermann et al. 2000; Caubit et al. 2004; Al-Mansoori et al. 2009; Amaechi et al. 2014). However, the detailed pore-scale (micrometre scale) fluid dynamics was poorly understood despite that pore-scale flow determines the overall reservoir-scale (kilometre scale) flow, in addition to other larger scale mechanisms, e.g. reservoir heterogeneity or gravity segregation (Bear, 1988). Therefore, it is vital to deepen our understanding in this area to aid the development of better storage mechanism techniques of CCS and EOR schemes, and to improve clean-up technologies for contaminated soils (see figures 2-4 and 2-5 below).



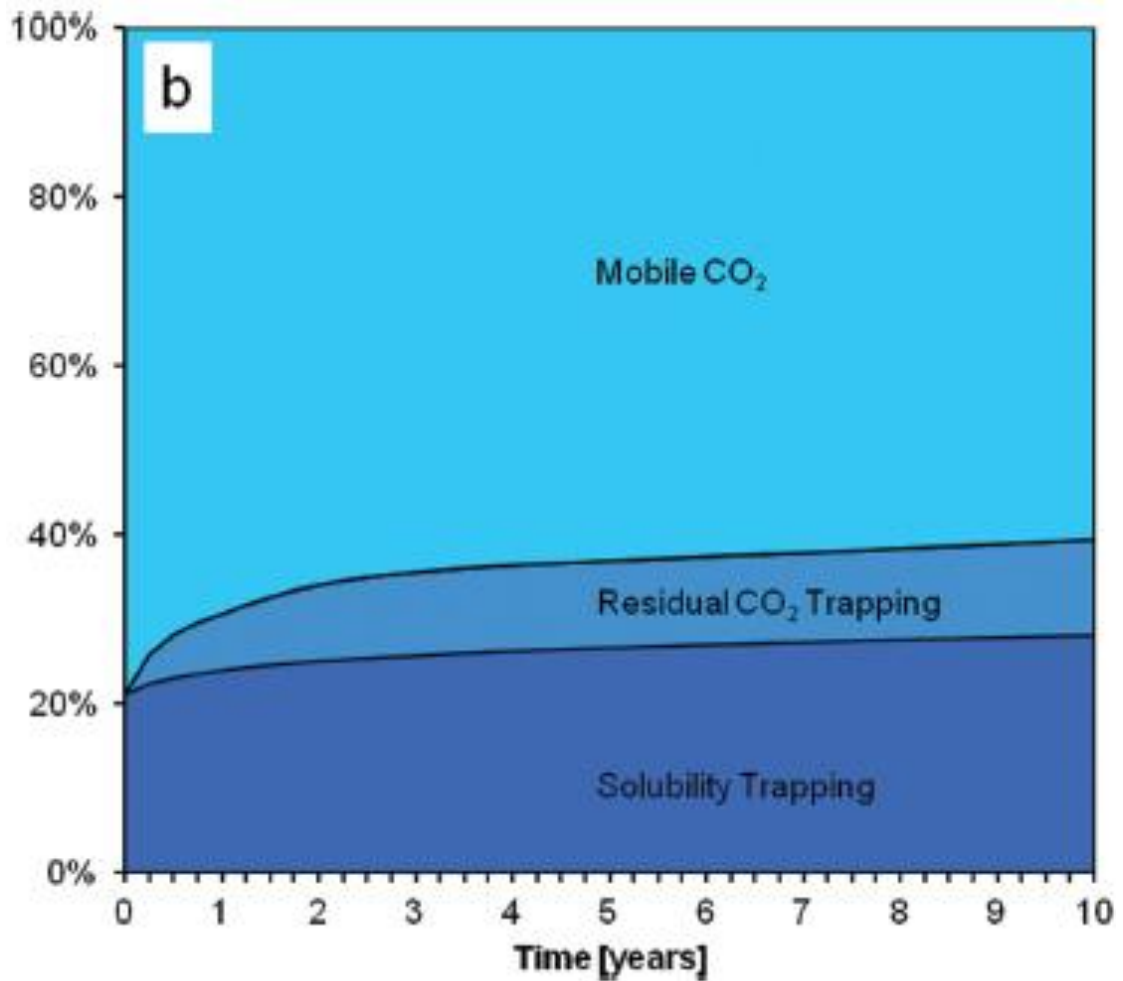


Figure 2-4 CO₂ fate simulation results from the modified tenth SPE Comparative Solution Project model using saturation functions derived from the contact angle to simulation workflow illustrated in Figure 2-5. (a) In the water-wet scenario, significantly more CO₂ is immobilised as a residually trapped phase than (b) in the CO₂-wet scenario, leading to significantly reduced CO₂ leakage risk. This figure was taken from Iglauer et al. (2015).

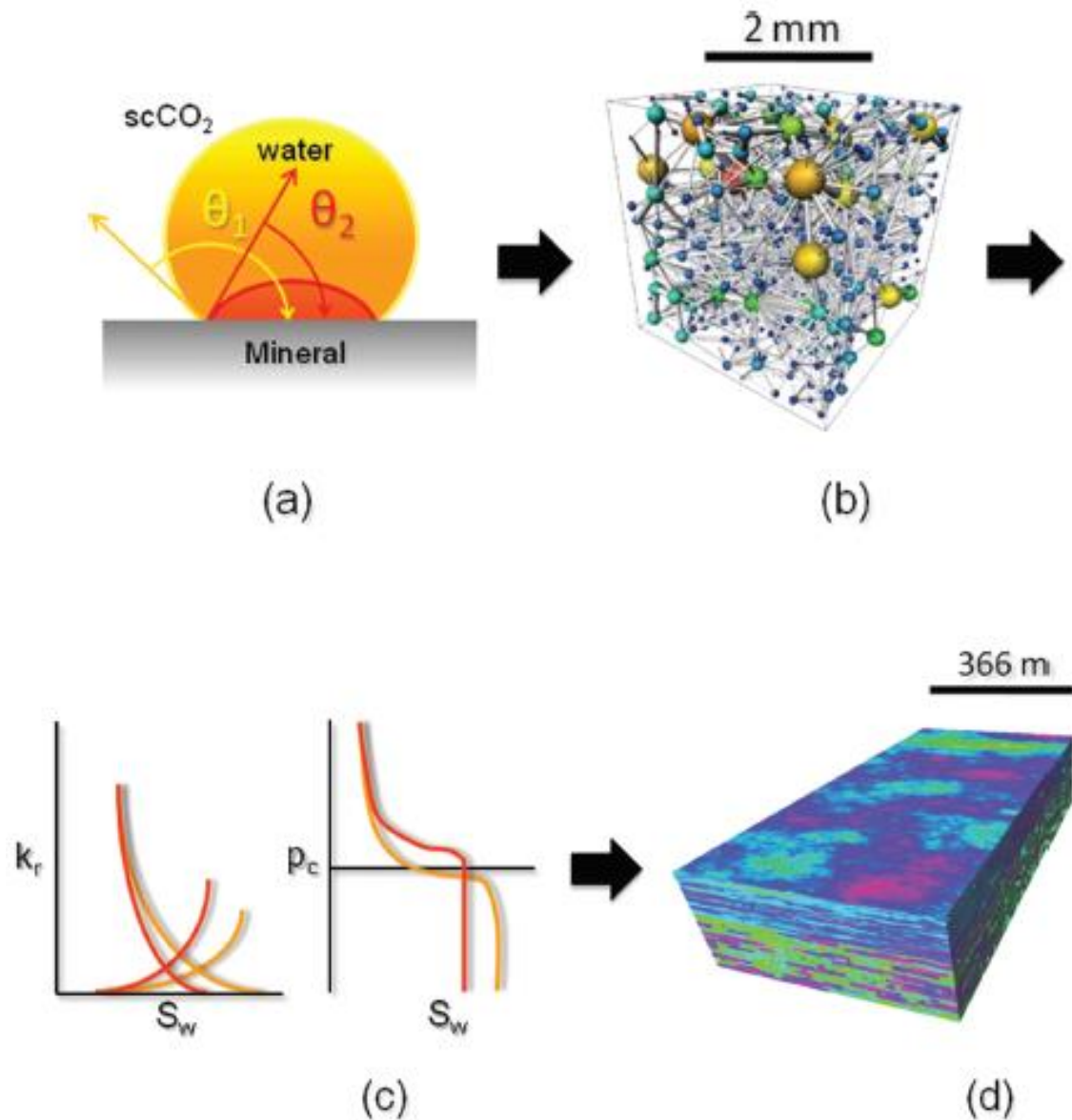


Figure 2-5 An example workflow leading from (a) millimetre-scale contact angle measurements to (d) hectometre-scale pilot-project simulation via (b) millimetre-scale pore-network modelling and (c) the predicted macroscale saturation functions for relative permeability and capillary pressure. The image of a pore network model was reproduced from Blunt et al. (2013). The permeability array is displayed per grid-block on a logarithmic scale in the pilot-scale model (d): the permeability scale is from 0.001 mD (green) to 10,000 mD (red) (9.87×10^{-19} - 9.87×10^{-12} m²). This figure was taken from Iglauer et al. (2015).

2.3 Dolomite Wettability after CO₂ injection

Wettability of a rock–CO₂–water system is a key parameter for CO₂ geo-sequestration and enhancing oil recovery owing to its impact on the three most important trapping mechanisms during the initial storage phase, i.e. structural (Naylor et al. 2011; Iglauer et al. 2015), residual (Iglauer et al. 2011, 2016; Chaudhary et al. 2013; Krevor et al. 2015; Al-Khdheawi et al. 2016, 2017, 2017a; Rahman et al. 2016) and dissolution trapping (Al-Khdheawi et al. 2016, 2017, 2017a).

Rock–CO₂–water wettability studies have been conducted by several researchers. For example, Pokrovsky et al. (Pokrovsky et al. 2005) observed there was no effect of salt concentration (0.1 to 1 M NaCl) on dolomite dissolution rate for CO₂ pressures ranging from 5 to 50 atm at 25°C and pH varying between 3 to 4, e.g. the reservoir pH values were between 3-4 at reservoir conditions for CO₂-saturated (“live”) brine (Schaefer et al. 2004). Moreover, Pokrovsky et al. (Pokrovsky et al. 2005) observed an insignificant effect of CO₂ on the rate of dissolution of calcite, dolomite, and magnesite where the dissolution rate stays constant when CO₂ pressure was further increased. Note that the apparent activation energy for calcite dissolution at 25–100 °C is equal to $48.2 \pm 4.6 \text{ kJ mol}^{-1}$ and 34 kJ mol^{-1} at pH equal to 4 (Pokrovsky et al. 2009). Therefore, CO₂ can be trapped for several hundreds of years owing to the slow dissolution kinetic triggered by the limited mixing between the CO₂ and the brine (Ballentine et al. 2001). Moreover, high temperatures (100 to 150 °C) and pressures (> 50 atm) have a weak effect on dissolution rates of all carbonate minerals (Ballentine et al. 2001), which enhances carbon dioxide sequestration in deep carbonate sedimentary basins.

The contact angle between the mineral-water-CO₂ was measured to quantify the wettability of the rock substrate (Iglauer et al. 2015a, 2017). It was reported that the water contact angle (θ) increased with increasing pressure owing to the reduction of the density difference between the fluids (Al-Yaseri et al. 2016; Rahman et al. 2016; Iglauer et al. 2017) and an associated increase in intermolecular interaction between CO₂ and minerals. This has been well-supported by most literature (Dietrich et al. 1991; Saraji et al. 2013; Iglauer et al. 2015a; Sarmadivaleh et al. 2015; Arif et al.

2016) and molecular dynamics simulations (Iglauer et al. 2012; Jung et al. 2012; McCaughan et al. 2013; Sedghi et al. 2014; Chen et al. 2015; Javanbakht et al. 2015).

Further, it was found that the contact angle increases with rising temperature (Yang et al. 2008; Farokhpoor et al. 2013; Iglauer et al. 2015; Sarmadivaleh et al. 2015; Al-Yaseri et al. 2016, 2016a; Roshan et al. 2016) but other researchers believe it is reduced (De Ruijter et al. 1998; Saraji et al. 2014; Arif et al. 2016). However, it was found that the contact angle is a complex function that includes temperature and many other parameters (Roshan et al. 2016).

Chapter 3 Effect of Temperature on CO₂/Brine/Dolomite Wettability: Hydrophilic vs Hydrophobic Surfaces*

Abstract

The water contact angle in a system of brine (20wt% CaCl₂) and CO₂ was measured on a smooth dolomite surface (RMS surface roughness 45nm) with both hydrophilic and hydrophobic behaviour as a function of pressure (0.1, 5, 10, 15, 20 MPa) and temperature (308, 323, and 343 K). The experimental results show that the contact angle of brine/CO₂ increases slightly with temperature when dolomite surface is hydrophilic but interestingly reduces when the surface is hydrophobic. The results also illustrate that the brine/CO₂ contact angles increase with increasing pressure. We interpreted the experimental observations using the concept of alteration in Van der Waals potential (substrate surface chemistry) with thermodynamics properties including pressure and temperature.

Keywords: Wettability, Dolomite, Contact angle, Carbon dioxide, Hydrophilic, Hydrophobic

*Reference: Al-Yaseri et al. (2017) *Journal of Energy and Fuel*, 31, 6329 - 6333

3.1 Introduction

CO₂ injection into carbonate reservoirs is a process for enhancing oil and gas recovery (Olenick et al. 1993; Babadagli 2006). The technique has been also suggested for CO₂ geo-sequestration (Babadagli, 2006; Izgec et al. 2008; Mohamed et al. 2011) which will be an efficient method to reduce anthropogenic CO₂ emissions to the atmosphere and thus mitigates climate change (Benson et al. 2005; Metz et al. 2005). However, it is important to understand the nature of interaction between fluid and rock in order to safely store CO₂ in the ground and assess the CO₂ ability to increase the recovery (Oomole et al. 1993; Iglauer et al. 2015, 2015a). Furthermore, the wettability of a rock–CO₂–water system is a key parameter in such process due to its impact on the three most important trapping mechanisms during the initial storage phase, i.e. structural (Naylor et al, 2011, Iglauer et al. 2015), residual (Iglauer et al. 2011; 2016, Krevor et al. 2015; Chaudhary et al. 2013; Rahman et al. 2016; Al-Khdheawi et al. 2016, 2017, 2017a) and dissolution trapping (Al-Khdheawi et al. 2016, 2017, 2017a).

Pokrovsky (Pokrovsky et al. 2005) observed that there was no effect of salt concentration (0.1 to 1 M NaCl) on dolomite dissolution rate for CO₂ pressures ranging from 5 to 50 atm at 25°C and pH varying between 3 to 4 e.g. the reservoir pH values are between 3-4 at reservoir conditions for CO₂-saturated (“live”) brine (Schaefer et al. 2004). Moreover, Pokrovsky et al. (Pokrovsky et al. 2005) observed insignificant effect of CO₂ on the rate of dissolution of calcite, dolomite and magnesite where the dissolution rate stayed constant when CO₂ pressure was further increased. Note that the apparent activation energy for calcite dissolution at 25–100 °C is equal to $48.2 \pm 4.6 \text{ kJ mol}^{-1}$ and 34 kJ mol^{-1} at pH equal to 4 (Pokrovsky et al. 2009). Therefore, CO₂ can be trapped for several hundreds of years due to the slow dissolution kinetic triggered by the limited mixing between the CO₂ and the brine (Ballentine et al. 2001). Moreover, high temperatures (100 to 150 °C) and pressures (> 50 atm) have very weak effect on dissolution rates of all carbonate minerals (Iglauer 2017) which enhances carbon dioxide sequestration in deep carbonate sedimentary basins.

In this context, the contact angle between the mineral-water-CO₂ was measured to quantify the wettability of the rock substrate (Iglauer et al. 2015a, 2017). It was reported that the water contact angle (θ) increases with increasing pressure due to the reduction of the density difference between the fluids (Al-Yaseri et al. 2016, Rahman et al. 2016, Iglauer 2017), and an associated increase in intermolecular interaction between CO₂ and mineral. This has been well-supported by most literature data (Dietrich et al. 1991, Saraji et al. 2013, Iglauer et al. 2015, Sarmadivaleh et al. 2015, Arif et al. 2016) and molecular dynamics simulations (Iglauer et al. 2012, Jung et al. 2012, McCaughan et al. 2013, Sedghi et al. 2014, Chen et al. 2015, Javanbakht et al. 2015).

The influence of temperature on contact angle has, however, stayed controversial with some believing that it increases with temperature; (Yang et al. 2008, Farokhpoor et al. 2013, Iglauer et al. 2015, Sarmadivaleh et al. 2015, Al-Yaseri et al. 2016, 2016a, Roshan et al. 2016), and others who argue that it reduces (De Ruijter et al. 1998, Love et al. 2005, Arif et al. 2016,). Roshan et al. (Roshan et al. 2016) showed theoretically and experimentally that the contact angle is a complex function of temperature through several parameters on electrically charged surfaces. However, a comprehensive quantitative analysis especially for weakly charged surfaces is required to give a conclusion to temperature dependency of contact angle.

Therefore, in this work the advancing and receding contact angles of brine (20 wt% CaCl₂)/CO₂ were measured on both, hydrophilic and hydrophobic, dolomite surfaces at different pressures (0.1, 5, 10, 15, 20 MPa) and temperatures (308, 323, and 343K). The results were analysed by extending the newly developed physical model for contact angle prediction (Al-Yaseri et al. 2016).

3.2 Experimental methodology

A dolomite sample with smooth surface (RMS roughness~110 nm measured with an atomic force microscopy instrument, model DSE 95-200, Figure 3-1) was used for the experiment. In order to measure the mineral composition of the dolomite sample, a sub-sample was first crushed in a tungsten carbide mill and the powdered sample

was mounted in an XRD sample holder and analysed using a PANalytical Empyrean II X-ray Diffractometer (45kV, 40mA; 5°-80° 2 θ), equipped with a Co anode source. The patterns were analysed and quantified using SiroQuant (v.4), with minerals identified with reference to ICDD powder diffraction files, Table 3-1.

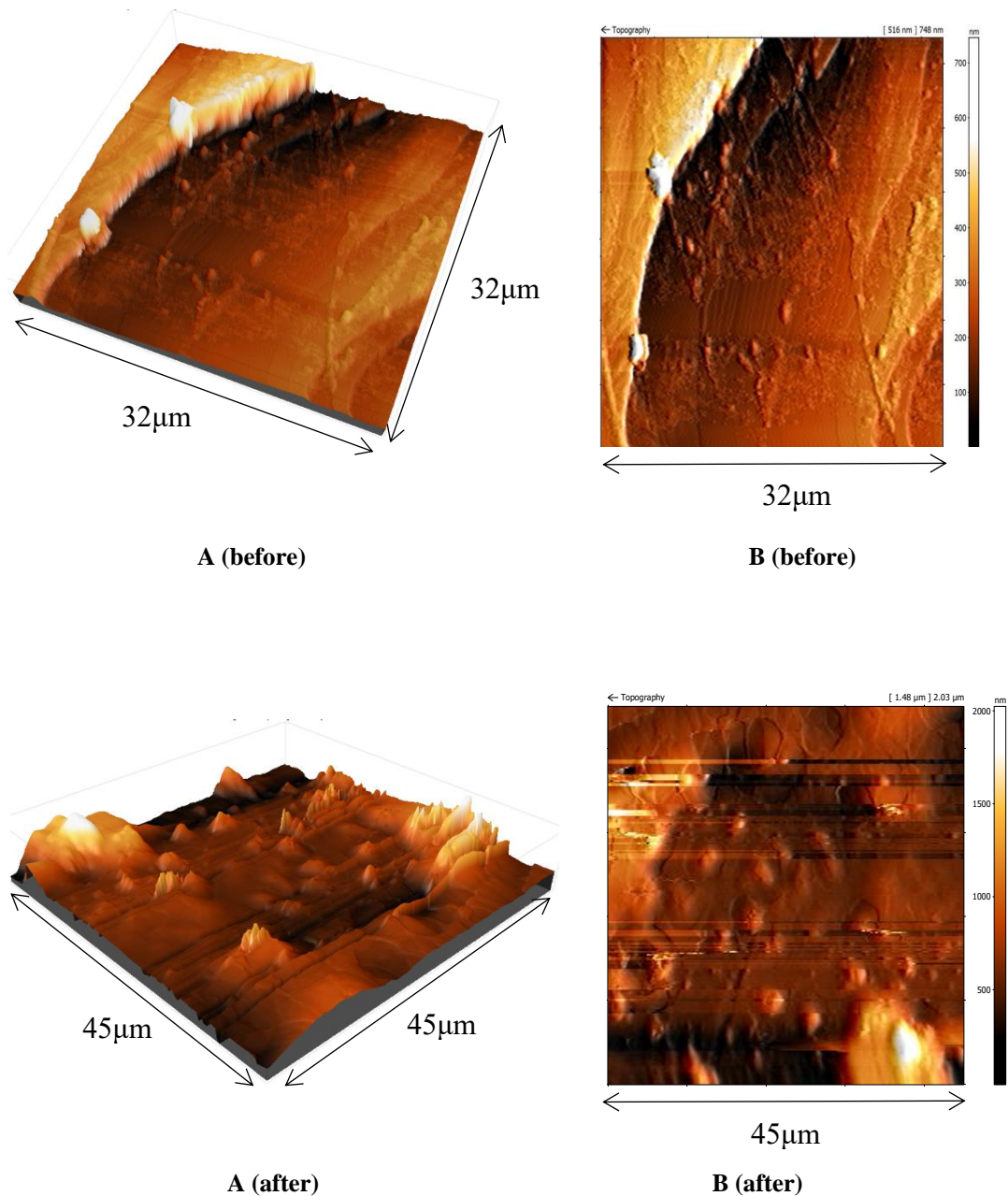


Figure 3-1 Atomic force microscopy images of the dolomite surface used in the experiments. (a) 3D topography of the substrate; (b) deflection signal, different heights are coloured differently (black is 0 nm, white is the maximum). For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.

Furthermore, the surface charge of the tested crushed dolomite sample suspended in brine (20 wt% CaCl₂) was measured (-6.24 mV) with a zeta potentiometer (Brookhaven ZetaPALS) at ambient conditions (298 K and 0.1 MPa).

Before each contact angle measurement, the dolomite (hydrophilic) sample was cleaned with acetone and DI water and then exposed to air plasma for 15 min to remove surface contaminations (Love et al. 2005, Mahadevan 2012, Iglauer et al. 2014). The substrate was then placed in a high pressure and temperature cell. The cell was pressurized with CO₂ using a high precision syringe pump (ISCO 500D; pressure accuracy of 0.1% FS) until a pre-set value was obtained. When the desired pressure was attained, a small droplet (~ 7 μ L) of brine (20 wt% CaCl₂ saturated by CO₂ using a mixing reactor; El-Maghraby et al. 2012) was dispensed onto the tilted substrate (tilting angle 12°) and advancing and receding contact angles were measured simultaneously using a high definition video camera. This procedure was repeated for different pressures (0.1, 5, 10, 15, and 20 MPa) and temperatures (308, 323, and 343 K).

The dolomite sample surface was initially hydrophilic as seen from measured contact angles. In order to make it hydrophobic for subsequent experiments, the sample was aged in Triethoxy(octyl) (99.9 mol% purity, from Sigma–Aldrich) silane for one month at ambient conditions and then left to dry at room temperature (Grate et al. 2012), note that the hydrophobic dolomite sample was cleaned by DI water only before each measurement.

Note that the surface roughness was measured by AFM before and after each contact angle measurements to assess the surface morphology and therefore possible dissolution-precipitation (Butt et al. 2006). The result showed insignificant change on the surface roughness (from 110 nm to 180 nm) which is unlikely to affect the contact angle values (Al-Yaseri et al. 2016).

Table 3-1 Minerals identified in Dolomite sample based on XRD pattern interpretation.

Phase	Weight	Error of Fit
Dolomite (CaMg(CO ₃) ₂)	96.9	1.25
Talc, (Mg ₃ Si ₄ O ₁₀ (OH) ₂)	3.1	1.25

3.3 Theoretical background

Young's equation (Young 1805) is often used to analyse the surface free energy of surfaces with no potential charges through contact angle measurement e.g. where the contact angle of liquid (θ) is related to the gas-solid (γ_{gs}), gas-liquid (γ_{lg}), and liquid-solid (γ_{ls}) interfacial tensions:

$$\cos \theta = \frac{(\gamma_{sg} - \gamma_{sl})}{\gamma_{lg}} \quad (1)$$

Difficulties of measuring the gas-solid and liquid-solid interfacial tensions, however, limit the applicability of Young's equation. The sharp-kink approximation (Dietrich et al. 1991, Merath, 2008) was therefore introduced to obtain an expression for contact angle without the need of gas-solid and liquid-solid interfacial tensions (Roshan et al. 2016):

$$\cos \theta = \frac{\Delta\rho}{\gamma_{lg}} I - 1 \quad (2)$$

Where $I = -\int_{z_{\min}}^{\infty} V(z)dz$ is the van der Waals potential integral (I) (Gatica et al. 2004, Garcia et al. 2008), and $\Delta\rho = \rho_{lf} - \rho_g$ (ρ_{lf} and ρ_g are densities of liquid film and gas respectively) (Merath, 2008).

It is evident from Eq. (2) that medium pressure influences the density difference and gas-liquid interfacial tension, however, the effect of pressure on density difference is much higher than its effect on γ_{lg} (Georgiadis et al. 2010). Increase in pressure significantly reduces $\Delta\rho$ and to some extent γ_{lg} .

The $\frac{\Delta\rho}{\gamma_{lg}}$ ratio therefore reduces with pressure thus increasing contact angles on any surfaces. In addition, the pressure effect is more significant than temperature effect on $\Delta\rho$. This is because ρ_g is a strong function of pressure and weaker function of temperature (Eq. (2)) e.g. especially for temperatures and pressures typically found in subsurface systems.

Furthermore, $\Delta\rho$, γ_{lg} and I are temperature dependent to some extent. In particular, temperature increase increases the gas-liquid interfacial tension and $\Delta\rho$ (especially at relatively higher pressure $\geq 5\text{MPa}$) however the $\Delta\rho$ increment with temperature is more significant than that of γ_{lg} (Sarmadivaleh et al. 2015, Georgiadis et al. 2010). This in turn means that $\frac{\Delta\rho}{\gamma_{lg}}$ increases with increasing temperature, which should thus reduce the contact angle e.g. assuming temperature independency of the van der Waals potentials. However, a contact angle increase with temperature has been observed experimentally.

Interestingly, it has been shown that the van der Waals potentials can decrease with temperature (Pinon et al. 2016) which can in turn increase the contact angle (Eq. (2)). The van der Waals potential ($V = V_s - V_l$) in Eq. (2) describes the net preference of the adsorbate molecule for wetting the substrate instead of forming a droplet, due to intermolecular forces (Garcia et al. 2008). In the definition of V , V_s is the potential energy of the adsorbate molecule due to the substrate and V_l is the potential energy of the adsorbate molecule due to the substrate hypothetically on the same location.

Thus, the temperature influences the van der Waals potential in both, hydrophilic and hydrophobic, surfaces however the hydrophilic surfaces will be more affected by

temperature. This is because hydrophilic surfaces offer more effective van der Waals interactions between the water molecules and substrate surface that can be broken by temperature increase (Pinon et al., 2016).

3.4 Results and discussion

The advancing and receding brine contact angles measured in the presence of CO₂ on hydrophilic dolomite sample at different temperatures (308, 323 and 343 K) and pressures (0.1, 5, 10, 15 and 20 MPa) are presented in Figure 3-2, respectively. An increase in pressure significantly increased both, the advancing and receding, contacts angles; however, a temperature increase either slightly increases the contact angle or in some cases has no effect on contact angles on the hydrophilic dolomite surface. However, on the hydrophobic dolomite sample, the temperature increase reduced the contact angle (Figure 3-2), while increasing pressure still increased the contact angle. Recall from section 3 pressure significantly influences the fluid densities and reduces the density difference between the fluids (Eq. (2)). This in turn increases the contact angle considerably in both hydrophilic and hydrophobic, surfaces.

The effect of temperature, on the other hand, is different for hydrophilic and hydrophobic dolomite surfaces. This difference was discussed earlier where the ratio

$\frac{\Delta\rho}{\gamma_{lg}}$ of γ_{lg} increases with temperature and therefore θ should in fact reduce on any surface. It seems to be the case for hydrophobic surfaces where the alteration of van der Waals potentials (I) with temperature is less significant than that of hydrophilic surfaces (Pinon et al. 2006).

Temperature can however reduce the van der Waals potential on hydrophilic surfaces more significantly. This is due to the fact that the hydrophilic surfaces offer more effective van der Waals interactions between the water molecules and substrate surface and therefore more potential energy. Such effective potential energy on hydrophilic surfaces is a strong function of temperature e.g. bound rupture can occur

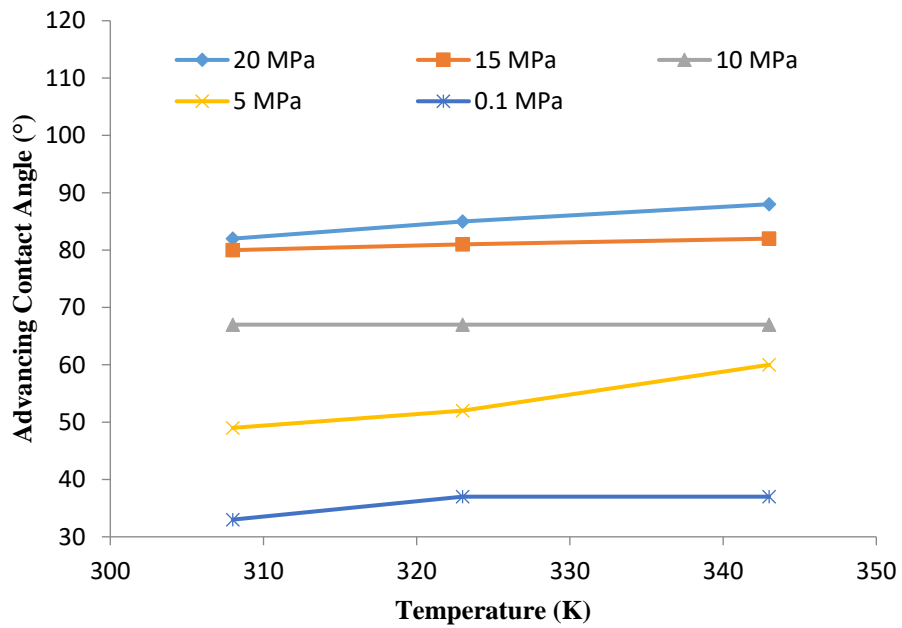
readily when temperature increases (Pinon et al. 2006). On hydrophobic surfaces, however such bounds are weakly present and therefore their temperature dependency is also weaker than hydrophilic surfaces.

Reduction in van der Waals potential increases the contact angle, thus two terms in

Eq. (2) ($\frac{\Delta\rho}{\gamma_{lg}}$ and I) compete with each other to define the contact angle. Therefore, the decrease in contact angle by temperature is more likely to be experienced in hydrophobic surfaces than hydrophilic. That is why the contact angle measured on hydrophobic surface (Fig. 6-3) decreases significantly with temperature when compare to hydrophilic surface (Fig. 6-2).

The above-mentioned behaviour was also seen by Takei et al. (Takei et al. 1994) in the field of bio-medical science. They used different type of polymers (N-isopropylacrylamide class) to modify the substrate surface from hydrophilic to hydrophobic by changing the temperature to control the drug delivery. They observed that the contact angle of water on the substrate was quite constant below a critical temperature (when the sample was hydrophilic). As soon as the critical temperature was reached (when the sample turned hydrophobic), the contact angle reduced significantly. The hydrophilic behaviour of the surface at relatively low temperature was linked to the hydrogen bonding which was later broken (when temperature passes critical point).

(a)



(b)

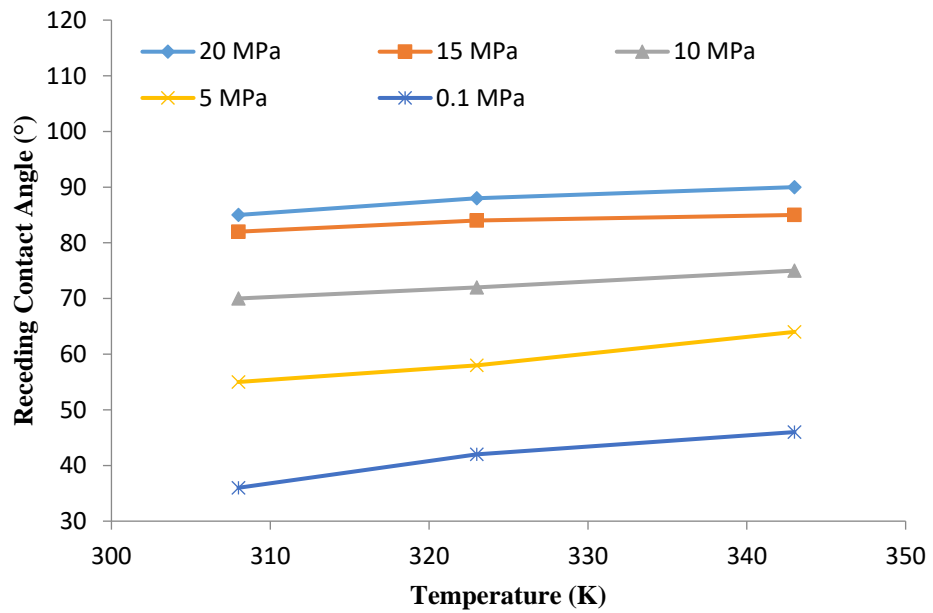
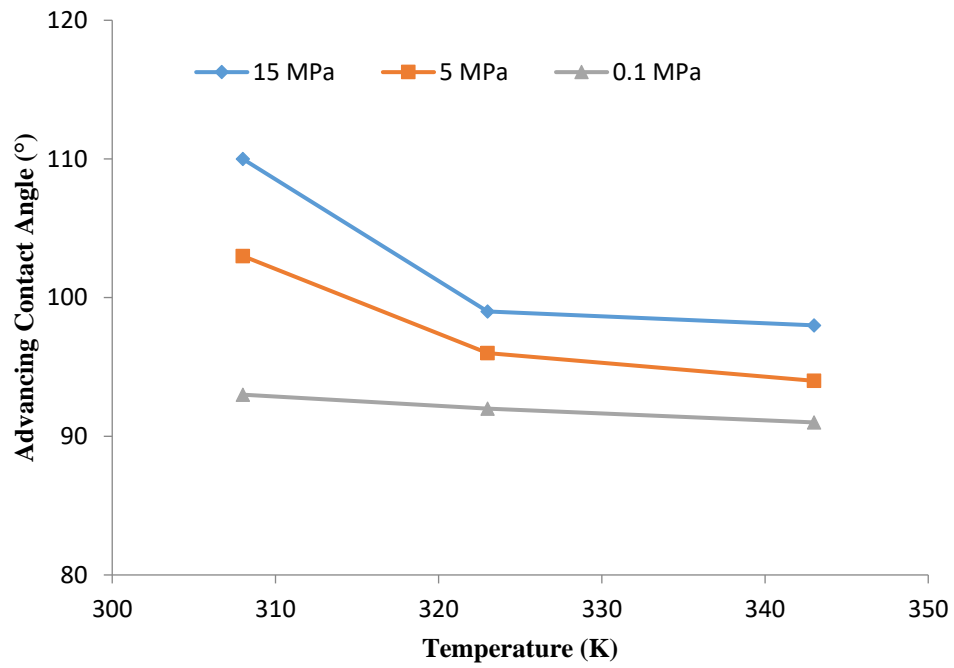


Figure 3-2 a) advancing and b) receding brine contact angles in the presence of CO₂ measured on hydrophilic dolomite at different pressures (0.1, 5, 10, 15 and 20 MPa) and temperatures (308, 323 and 343K).

(a)



(b)

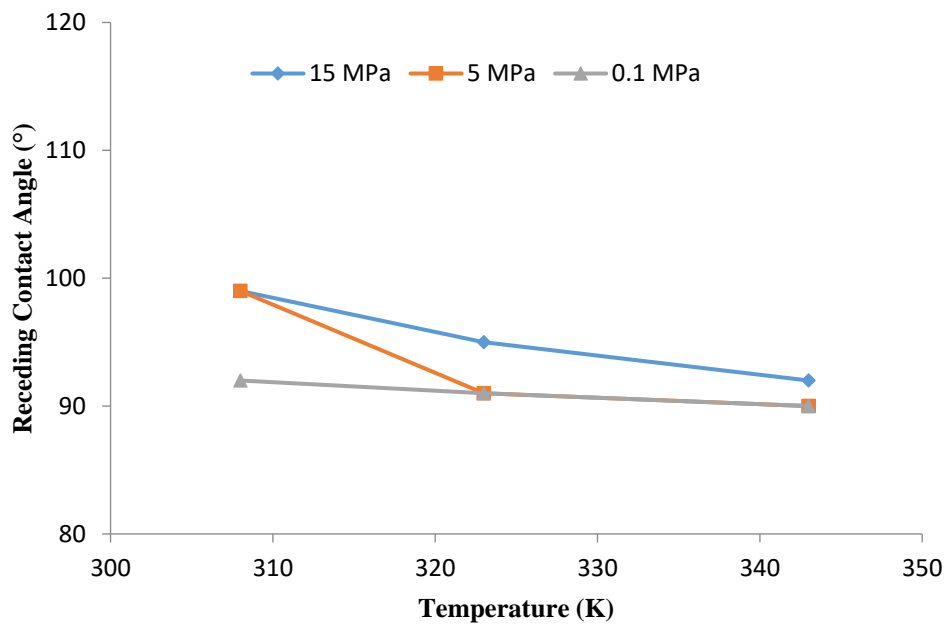


Figure 3-3 a) advancing and b) receding brine contact angles in the presence of CO₂ measured on hydrophobic dolomite at different pressures (0.1, 5 and 15 MPa) and temperatures (308, 323 and 343K).

3.5 Conclusions

In this chapter, we measured the advancing and receding contact angles of CO₂/brine system on hydrophobic and hydrophilic dolomite surfaces at different pressures and temperatures. We analysed the experimental observations with the concept of temperature-dependency of van der Waals potential and sharp-kink approximation. It was shown that pressure increases the contact angle of brine considerably regardless of surface being hydrophilic or hydrophobic. Contact angle on the other hand stayed unchanged or slightly increased on hydrophilic surfaces by increasing temperature but were reduced on hydrophobic surfaces. It was illustrated that phase-density difference, fluids interfacial tension and fluids-solid van der Waals potentials affect the contact angle of no/weakly charged surfaces when state variables such as pressure or temperature vary.

Chapter 4 Residual trapping of supercritical CO₂: direct pore-scale observation using a low-cost pressure cell for micro-computer tomography*

Abstract

We describe the design and fabrication steps of a low-cost, modular experimental cell for flooding and X-ray tomography visualization of the rock samples at reservoir conditions (confining pressure up to 45 MPa, pore fluid pressure up to 20 MPa, and temperature up to 80°C). The main body of the cell is made from Polyether ether ketone (PEEK) tube. The cell can be easily machined and assembled in house from off-the-shelf components at a low cost. Modular design allows adopting the cell to a wide range of X-ray visualization experiments. To demonstrate the capabilities of the equipment we are presenting how the residual CO₂ phase spreads in the rock as a function of thermo-physical conditions, namely, for ambient condition CO₂ gas, near-critical CO₂ gas, liquid CO₂, and supercritical CO₂.

Keywords: CO₂, Rocks, Micro CT, Pressure cell

*Reference: Lebedev et al. (2017) in *Journal of Energy Procedia*, 114, 4967 – 4974

4.1 Introduction

Injection of CO₂ into the subsurface is one of the known methods to mitigate greenhouse gas emissions (Lackner 2003, Metz et al. 2005, Orr 2009). The buoyant CO₂ is contained in the subsurface by structural, residual, dissolution and mineral trapping mechanisms. Quantitative knowledge of the residual trapping capacity of the reservoir is thus crucial for storage site selection. This residual saturation depends on porosity, initial CO₂ saturation, pore morphology and wettability, or in other words on reservoir properties at the microscopic level; note that the macroscopic residual saturation also depends on sweep efficiency and gravity. Therefore, it is of crucial importance to understanding the major factors, which determine residual trapping efficiency. Micro-CT tomography technique is a powerful tool to understand such factors at micro-level.

It is standard practice in micro-CT to image rocks at ambient conditions (i.e. at room pressure and temperature) (Madonna et al. 2013, Shulakova et al. 2013). However, reservoir rocks are buried at such depth that they experience high stresses and temperatures. The thermodynamic properties of the fluids inside the reservoir are pressure and temperature dependent. Therefore, transport properties are also temperature and pressure dependent. Moreover, it is well established that elastic rock properties of rocks are strongly affected by stress and or fluid distribution. Thus, in order to acquire realistic pore network structures and fluid distributions (including but not limited to residual saturation), and to reliably estimate transport and elastic properties from micro images, rocks have to be imaged at reservoir pressure and temperature conditions. At this moment, cells for micro-CT are custom made (Iglauer et al. 2011, Fuisseis et al. 2014) and thus are expensive.

We thus have developed suit of X-ray transparent pressure and temperature (P-T) cells capable of holding from 5 mm to 38.5 mm diameter samples. To demonstrate capabilities of the equipment, in this paper we report how the residual CO₂ saturation depends on the thermodynamical conditions of the CO₂ (which changes from gas to liquid, near-supercritical and supercritical), in order to link CO₂ saturation and fluid distribution to acoustic properties of the samples.

4.2 Pressure cells for microCT

We have designed and fabricated two X-ray transparent pressure – temperature cell: both cells are utilized for visualization of the process inside the reservoir samples during flooding. The first cell is prepared for the simultaneous acoustic measurements and CT imaging at moderate resolution (40 μm) of the 38.5 mm (1.5”) in diameter samples; the second one is applied for micro-CT imaging of smaller samples (5mm in diameter) at sub-micrometer resolution. Cell for large samples is shown in Figure 4-1(a), this cell has been described in detail previously (Lopes et al. 2014, Lebedev et al. 2014).

Figure 4-1 (b) shows the cell for high-resolution imaging. This cell can be easily fabricated and assembled from off-the-shelf components at a low cost. Modular design allows adaptation of the cell to a wide range of X-ray experiments. The main body of the pressure cell is made from a 21-mm-diameter, 300-mm-long PEEK (Polyether ether ketone) rod supplied by RS components. An 8 mm in diameter hole to accommodate the sample inside was drilled in the laboratories’ workshop using a lathe. All other parts of the cell have been assembled from the standard Swagelok fittings. Both ends of the cell have the same assembly. Stainless steel, $\frac{3}{4}$ ” nuts with ferrule are swaged on both ends of the PEEK tube. The $\frac{3}{4}$ ” to $\frac{1}{4}$ ” unions are connected to both ends of the PEEK rod. The $\frac{1}{4}$ ” tees unions fittings are using for confining pressure supply and getting through of two $\frac{1}{8}$ ” stainless steel fluid lines inside the cell. The cell is fixed on the rigid holder that can be easily connected to a stage inside micro-computed tomography (μCT). A 5-mm-diameter sample was jacketed in the rubber sleeve and was connected to the fluid lines using custom made plastic plates. Some steps of the fabrication and assembly and testing of the cell are shown in Figure 4-2.

The arrangement of the cell inside X-ray micro-CT (3D X-ray Microscope VersaXRM 500 (XRadia-Zeiss)) is shown in the Figure 4-3. X-Y-Z micropositioning stage of the X-ray micro-CT is able to hold up to 10 kg weight; the weight of the developed cells is not exceed kg, thus the pressure cell with can be fixed to X-Y-Z stage without compromising performance of the scanner. All hydraulic and fluid lines of the cell are connected to equipment by flexible PEEK $\frac{1}{8}$ ” tubes (Sigma-

Aldrich). Such flexible tubes allow the rotation of the cell from -180° to $+180^{\circ}$ during the imaging. Confining pressure to a sample is applied by a set of syringe pumps (Teledyne-Isco 260 and 500 series). Depending on the requirements of the experiment, the heating and cooling of the sample inside the cell can be done by using continuous circulating of the pressurized confining fluid. A small flexible electrical heater attached to the low part of the cell (not shown in the Figures) can be used to boost the heating process. Injection of the fluids into the sample is performed using two syringe pumps (Teledyne-Isco 100 and 260 series). All pumps are placed in an outside X-ray micro-CT. All pumps and connecting tubes are maintained at a given temperature within the range -15°C $+80^{\circ}\text{C}$ by circulating of liquid antifreeze.

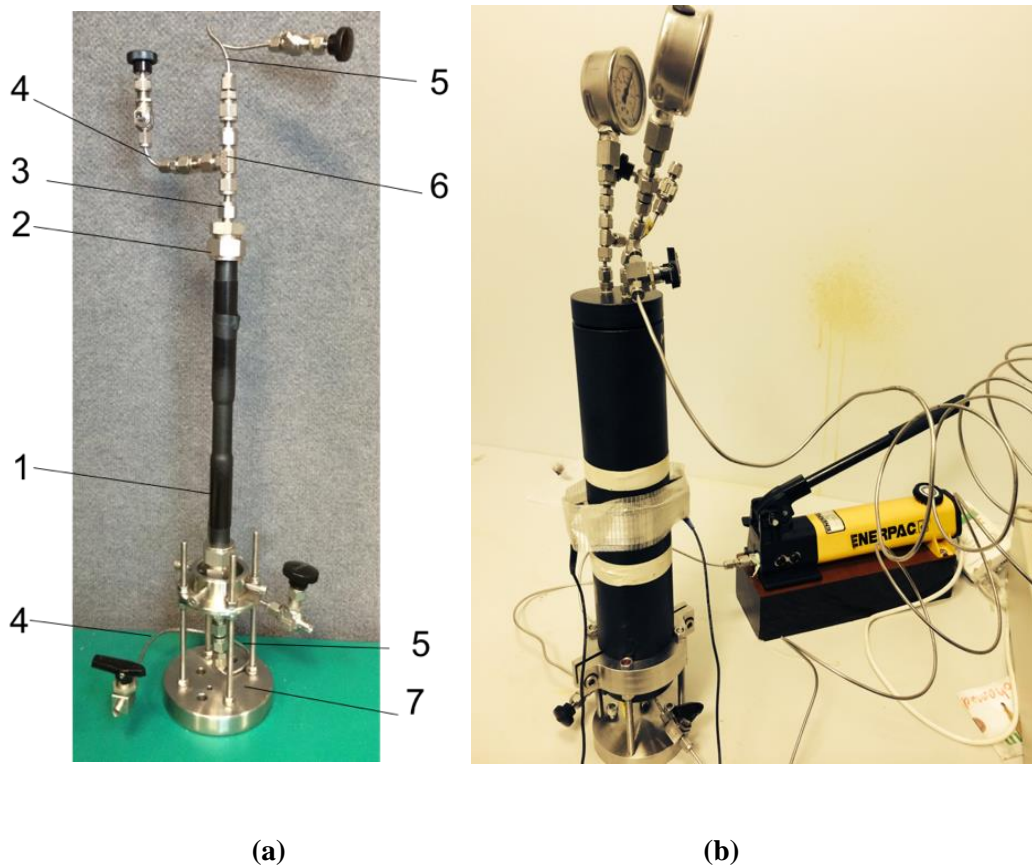


Figure 4-1 Pressure-Temperature cells for CT: a) pressure cell for 5 mm diameter samples; b) pressure cell for 38.5 mm diameter samples. 1) PEEK pipe; 2) Stainless steel, $\frac{3}{4}$ " nuts with ferrule; 3) $\frac{3}{4}$ " to $\frac{1}{4}$ " unions; 4) confining pressure line $\frac{1}{8}$ "; 5) pore pressure/ flooding line; 6) $\frac{1}{4}$ " tees union; 7) stage to be fixed to X-Y-Z stage of the micro-CT.

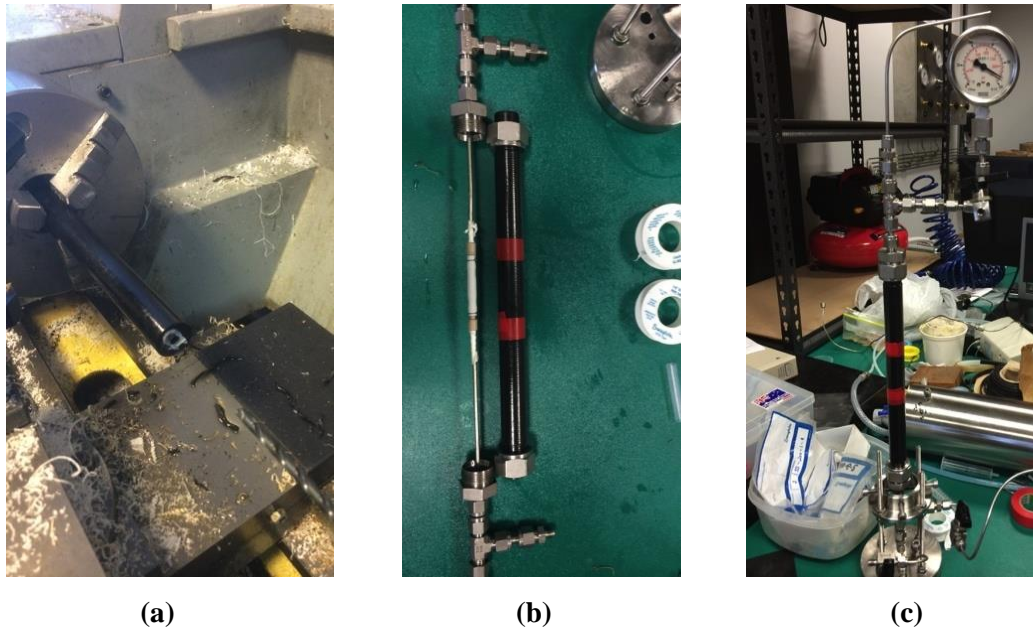


Figure 4-2 (a) Fabrication of the PEEK tube in the lab; (b) cell assembly: sample inside the white rubber sleeve is connected to flooding lines; (c) hydraulic test outside the micro-CT (confining pressure is 33 MPa, zero pore pressure).

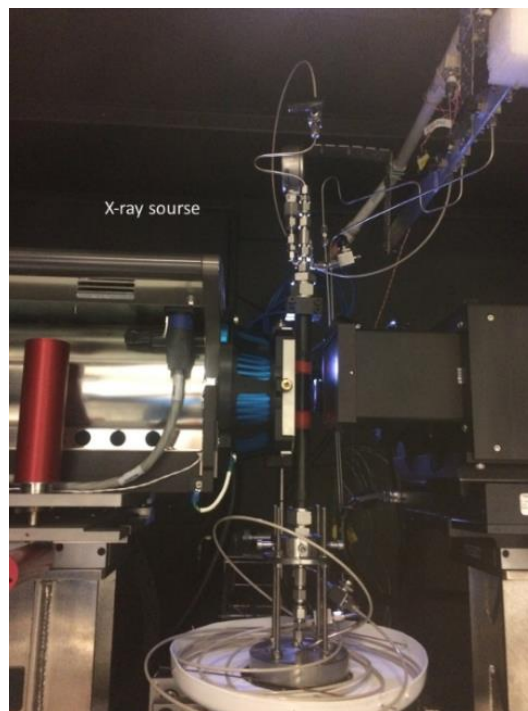


Figure 4-3 Pressure - Temperature cell inside micro-CY. X-Ray source in on the left, X-Ray objective (detector) is on the right.

4.3 Demonstration experiments

The cylindrical sample has been drilled from the bigger block, the sample is trimmed to size to accommodate inside a rubber sleeve inside the cell. After pressure test outside, the cell is placed inside micro-CT and the test image is performed. The cell and all hydraulic systems inside and outside the micro-CT are heated to designated temperature and controlled within 0.5C accuracy.

Monitoring and control of fluid injection inside the sample are performed between high-resolution imaging. Such real time in situ flooding visualization technique allows to estimate the amount of pore volumes (PV) injected inside the sample more precisely. Figure 4-4 shows the sequence of X-ray radiographs, i.e. one plane projection, of the Bentheimer sandstone sample inside the cell acquired at 20 MPa of confining pressure. Radiograph (a) is taken for the dry sample. On the radiograph (b) it can be observed that brine (doped with ten presents of sodium iodate) is entering the sample from the bottom and coming out the sample from the top. The last image (c) is the image after liquid CO₂ injection at 10 MPa injection pressure.

At each stage of the experiment, 3D imaging of the sample is performing at X-Ray energy ranged from 40kV to 80kV. Minimum time to acquire one 3D image is about 1.5h. During the acquisition, the injection valve is closed thus no fluid is injecting into the sample. However, the backpressure is maintained by a production pump. Figure 4-5 shows 2D images (sliced from 3D) of the dry, brine saturated and after liquid CO₂ flooding Bentheimer sandstone acquired at confining pressure of 20 MPa, pore pressure of 10 MPa, and temperature of 28°C with a nominal voxels size resolution of $(3.4348 \mu\text{m})^3$.

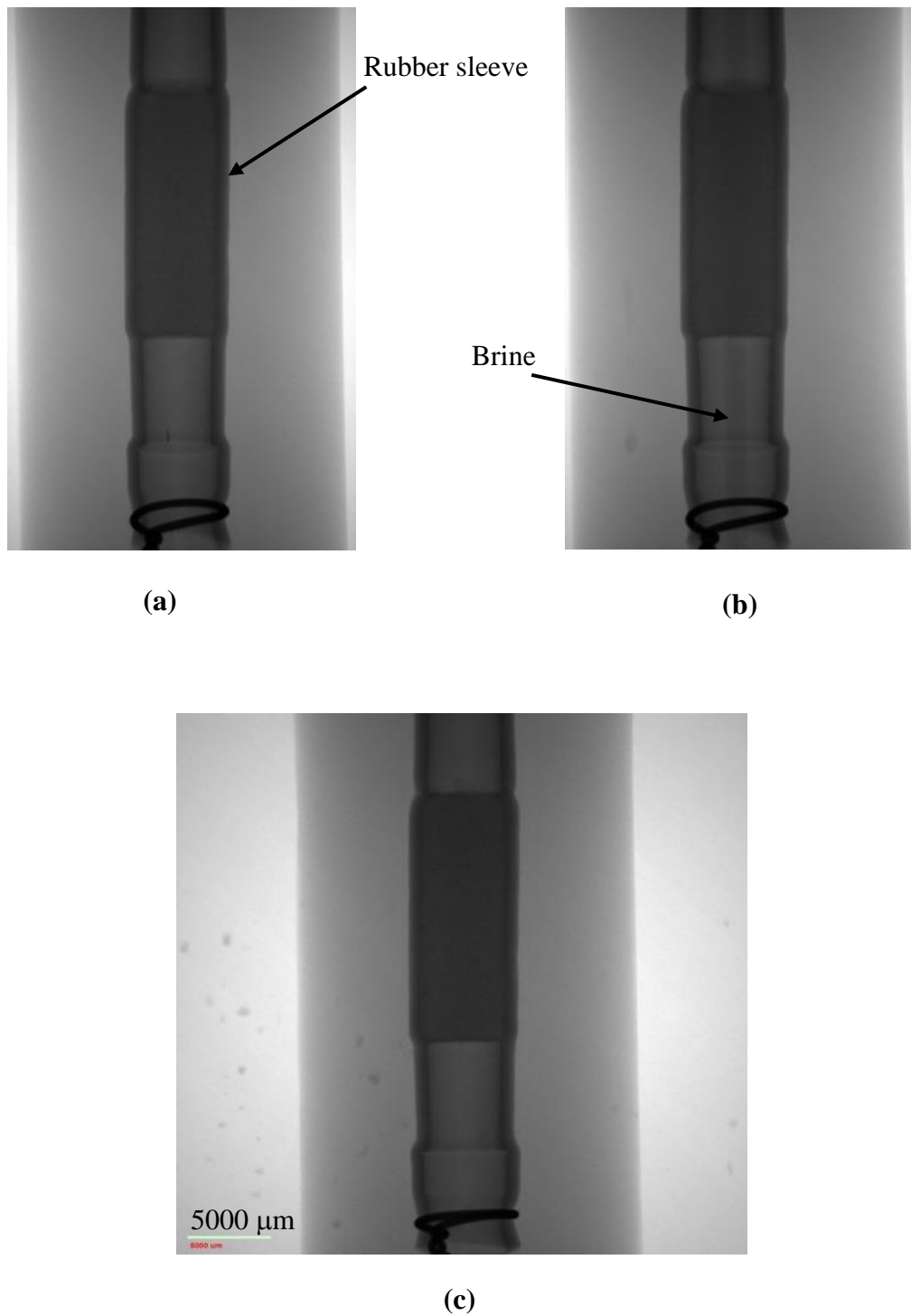
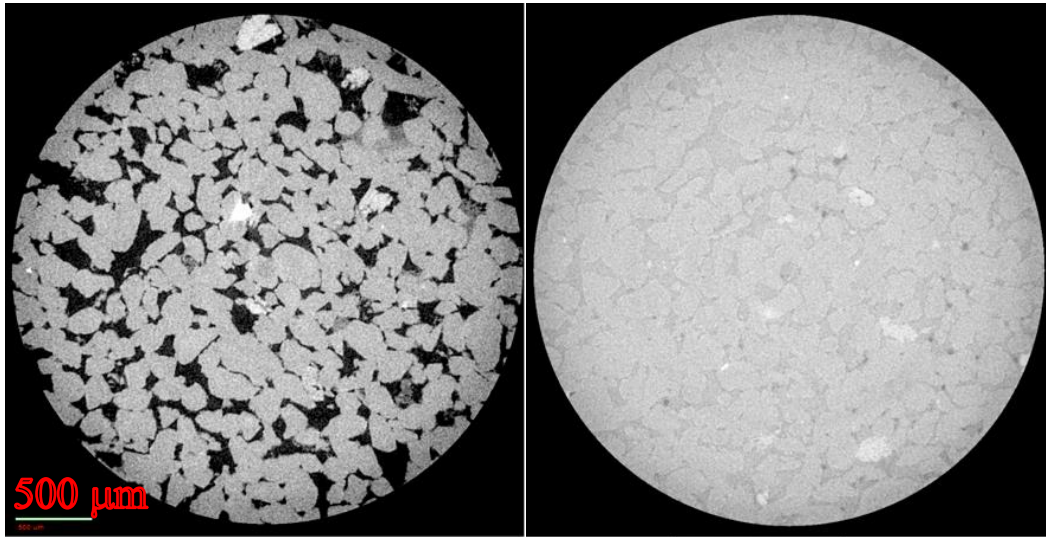
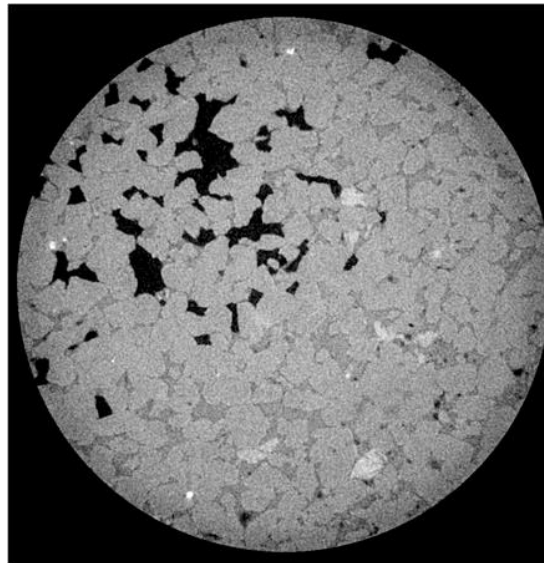


Figure 4-4 Time lapse of X-ray radiographs of the sample inside the cell at 20 MPa confining pressure: (a) dry sample; (b) brine injection; (c) after CO₂ injection. Low resolution image using 0.4x objective.



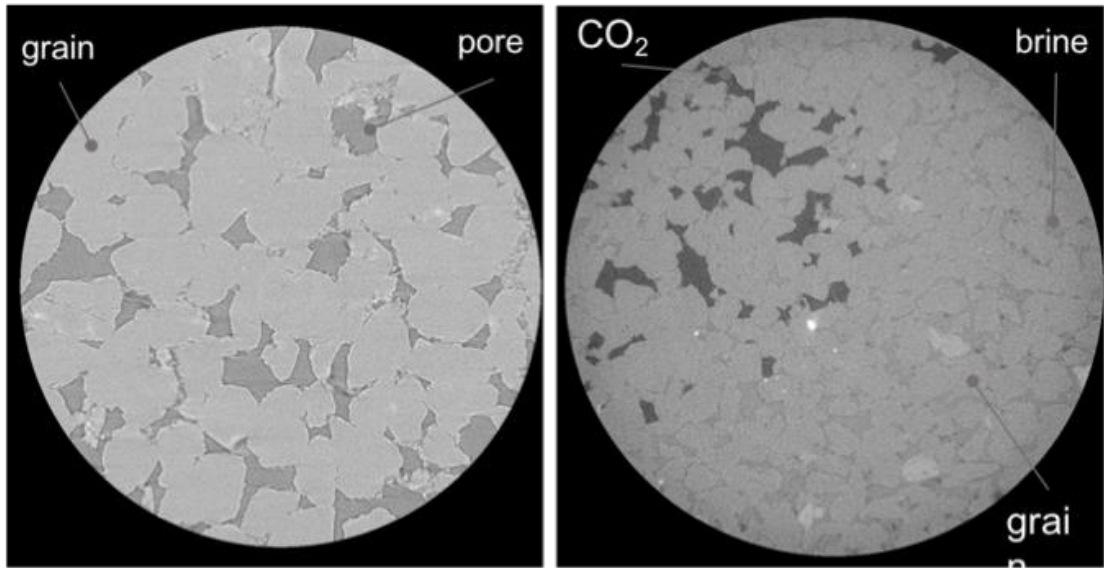
(a)

(b)



(c)

Figure 4-5 Tomographic images in the flooding experiment. (a) dry sample; (b) 10%NaI-doped brine saturated sample; (c) after flooding with liquid CO₂. The direction of CO₂ injection is perpendicular to the paper plane. X-ray energy is 80kV, the distance between X-Ray source and the sample is 20 mm, the distance between the sample and 4x objective is 20 mm; pixel size is 3.4348 micrometer. Confining pressure is 20 MPa, pore pressure is 10 MPa, the temperature is 28°C. Note that the images (b) and (c) are corresponding to the same area of the sample.

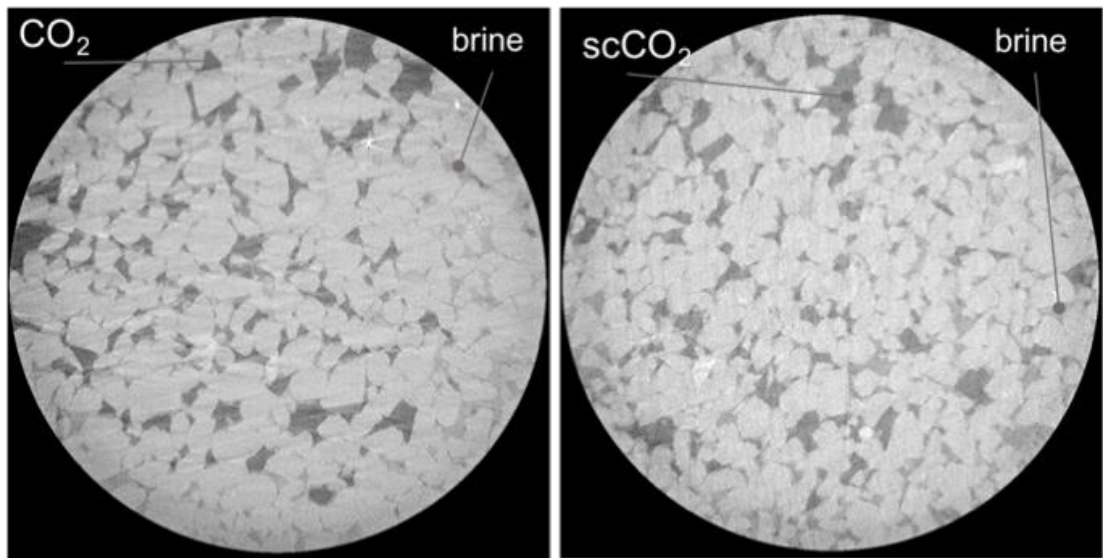


1.58 mm

3.43 mm

(a)

(b)



3.43 mm

3.43 mm

(c)

(d)

Figure 4-6 Influence of Pressure, Temperature and Fluid on the image quality. Bernheimer sandstone, gas permeability 1880mD. X-Ray energy is 60kV, the distance between X-Ray source and the sample is 19 mm, the distance between the sample and 4x objective is 19 mm; pixel size is 3.4348 micrometer. Number of scans (radiographs) is 4001: (a) Dry sample, $P_{conf}=30$ MPa, $T=28^{\circ}\text{C}$; (b) after liquid CO_2 injection into brine saturated sample, $P_{conf}=25$ MPa, $P_{inj}=10$ MPa, $T=28^{\circ}\text{C}$; (c) after “near critical” CO_2 injection into brine saturated sample $P_{conf}=15$ MPa, $P_{inj}=7$ MPa, $T=42^{\circ}\text{C}$; (d) after supercritical- CO_2 injection into brine saturated sample $P_{conf}=15$ MPa, $P_{inj}=10$ MPa, $T=42^{\circ}\text{C}$.

By increasing the pressure and especially temperature during the experiment the quality of raw images can decrease. Such slight degradation in quality caused mainly by the sample drift induced by a circulating high pressure and temperature confining fluid. Figure 4-6 shows how raw image quality changes when pressure and temperature increases.

4.4 Conclusion

The simple experimental cells that allow micro tomographic visualization of fluids inside rocks under elevated pressure was presented. The modular cell design allows hassle-free maintenance and easily modifications of the cell. The cell can be used for visualisation of multiphase fluids distribution inside porous media.

Using developed cell, we investigated how the residual CO_2 phase spreads in the rock as a function of thermo-physical conditions (for ambient condition CO_2 gas, liquid CO_2 , near-critical CO_2 gas, and supercritical CO_2); and how residual CO_2 is dissolved by under saturated formation brine (dissolution trapping). We found that CO_2 spread is relatively homogeneous throughout the core plugs for all conditions; however, less sc CO_2 was trapped at higher pressures, which is possible due to the lower water-wettability of sandstone at high pressures. We conclude that μCT is a very powerful tool for non-destructive visualisation of microstructure as well as pore-scale fluid distribution inside rocks samples. So far, the micro-CT cell has been successfully used in the following experiments pressure research (Saenger et al. 2016, Rahman et al. 2016, Zhang et al. 2016, 2016a, 2016b).

Chapter 5 Residual Trapping of Supercritical CO₂ in Oil-wet Sandstone*

Abstract

Residual trapping, a key CO₂ geo-storage mechanism during the first decades of a sequestration project, immobilizes micrometre sized CO₂ bubbles in the pore network of the rock. This mechanism has been proven to work in clean sandstones and carbonates; however, this mechanism has not been proven for the economically most important storage sites into which CO₂ will be initially injected at industrial scale, namely oil reservoirs. The key difference is that oil reservoirs are typically oil-wet or intermediate-wet, and it is clear that associated pore-scale capillary forces are different. And this difference in capillary forces clearly reduces the capillary trapping capacity (residual trapping) as we demonstrate here. For an oil-wet rock (water contact angle $\theta = 130^\circ$) residual CO₂ saturation $S_{CO_2,r}$ ($\approx 8\%$) was approximately halved when compared to a strongly water-wet rock ($\theta = 0^\circ$; $S_{CO_2,r} \approx 15\%$). Consequently, residual trapping is less efficient in oil-wet reservoirs.

Keywords: CO₂ geo-sequestration; Residual trapping; Storage capacity; Wettability; Oil-wet; Oil reservoir

*Reference: Rahman et al. (2016) in *Journal of Colloid and Interface Science*, 469, 63-68

5.1 Introduction

Carbon geo-sequestration has been identified as a feasible technology to mitigate global warming (Lackner 2003, Metz et al. 2005, Orr 2009). Technically, CO₂ is captured from large emitters (e.g. coal-fired power plants), and injected deep underground into geological formations for storage. However, although CO₂ is in a dense supercritical state at reservoir conditions (below 800m depth), it migrates upwards as it has a lower density than the resident formation brine. One key mechanism, which prevents the CO₂ from leaking back to the surface is residual trapping, where the CO₂ plume is split into many micrometre sized bubbles which are immobilized by capillary forces in the pore network of the rock (Suekane et al. 2006, Sukena et al. 2008, Pentland et al. 2011, Iglauer et al. 2011a, Krevor et al. 2012). Pore-scale residual trapping has been proven to work in clean sandstone (Iglauer et al. 2011a) and carbonate (Andrew et al. 2013, 2014). However, this mechanism has not been proven for oil-wet rock, despite its key importance as initial industrial scale CO₂ storage projects are very likely to occur in oil reservoirs (Firoozabadi et al. 2010); and these oil reservoirs are typically oil-wet (Cantucci et al. 2009). The significance of oil reservoirs for carbon storage is high, as sequestration can be directly combined with CO₂ driven enhanced oil recovery (Cuiec 1991, Emberley et al. 2005, Lake 2010); furthermore, depleted oil reservoirs already have required infrastructure in place (Cook 2014), and they are typically well characterised in terms of seismic surveys, which significantly aids (reservoir scale) CO₂ flow monitoring (Gurevich et al. 2014).

We thus imaged initial and residual CO₂ in an oil-wet rock at typical storage conditions with an x-ray micro-computed tomograph (μ CT), and we demonstrate that the residual trapping capacity is significantly reduced in oil-wet rock at (high pressure) storage conditions.

5.2 Experimental Procedure

5.2.1 Simulating Subsurface Condition

A sandstone oil reservoir at a depth of approximately 1000m was simulated (10MPa pore pressure, 318K temperature). For this, a clean homogeneous sandstone (Bentheimer) was selected to represent storage rock. Small core plugs (5mm diameter and 10mm length) were drilled and used for the experiments. Porosity was 22%, permeability 1800mD, and the rock consisted of 99wt% quartz, 0.7wt% kaolinite, and 0.3wt% rutile.

This rock was completely water-wet at ambient conditions (0° water contact angle in air), note that at higher pressures quartz is less water-wet (Chiquet et al. 2007, Broseta et al. 2012, Al-Yaseri et al. 2015). To mimic an oil reservoir the wettability of the rock was altered by immersing the plug into silane (Dodecyltriethoxysilane, 99.9mol% purity, from Sigma-Aldrich) under vacuum (so that the plug was fully saturated with the silane) and ageing the plug for 4 weeks at ambient conditions (Al-Anssari et al. 2016). This process rendered the plug strongly oil-wet (130° water contact angle at ambient conditions in air).

5.2.2 Imaging scCO₂ at the pore-scale with μ CT

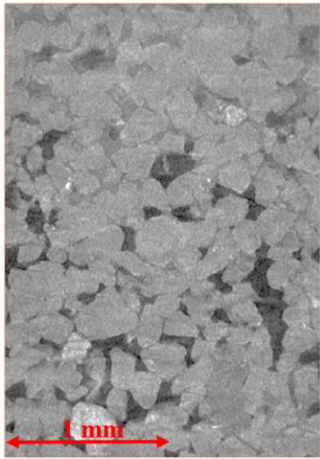
The oil-wet core plug was mounted into a high pressure μ CT cell (Iglauer et al. 2011a), vacuumed and fully saturated with doped “dead” brine (7wt% NaI in deionized water; note that NaI is required for CT contrast). Subsequently 15 MPa confining pressure was applied, and pore pressure was raised to 10MPa. All flow lines and fluids were isothermally heated to 318K. These conditions approximately correspond to a storage depth of 1000m, see above. The dead brine was then displaced with >40 pore volumes (PV) of “live” brine (live brine is brine thermodynamically equilibrated with CO₂; (El-Maghraby et al. 2012)). Then 40PV of supercritical (sc) CO₂ were injected at a flow rate of 0.1 mL/min, which corresponded to a capillary number of 6×10^{-8} , and the core plug was imaged in 3D at a resolution of $(3.4\mu\text{m})^3$ with an Xradia VersaXRM μ CT instrument. Finally, live brine (40PV) was injected at a flow rate of 0.1mL/min, which corresponded to a capillary number of 1.7×10^{-6} , and the sample was μ CT imaged again. This

experiment was repeated again, and the results were reproducible. The μ CT images were filtered with a 3D non-local means algorithm (Buades et al. 2005) and segmented with a watershed algorithm (Schlüter et al. 2014). Petrophysical properties were then measured on the images, see below. For comparison, the same experiment was conducted again with a water-wet Bentheimer plug (0° water contact angle, see above).

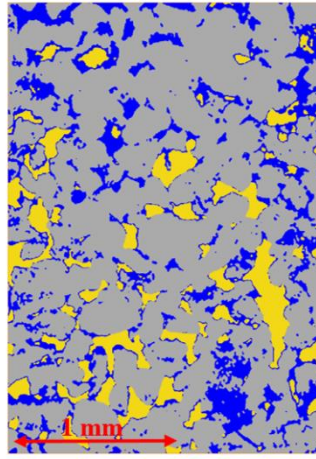
5.3 Results and Discussion

5.3.1 Visualization, cluster morphologies and initial & residual CO₂ Saturations

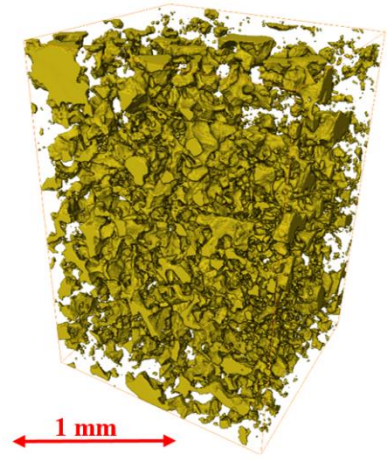
Initial and residual CO₂ saturations of the oil-wet and water-wet plugs are visualized in Figure 5-1. Initial CO₂ saturations ($S_{CO_2,i}$) were approximately the same ($\sim 50\%$), but the residual saturation ($S_{CO_2,r}$) in the oil-wet plug was significantly lower ($\approx 8\%$) than in the water-wet plug ($S_{CO_2,r} \approx 15\%$), Table 5-1. Precisely, the $S_{CO_2,r}$ measured for the water-wet plug was 14.9%, consistent with independent measurements (Sukena et al. 2008, Pentland et al. 2011, Iglauer et al. 2011a, Andrew et al. 2013, Krevor et al. 2015, Akbarabadi et al. 2013); note that $S_{CO_2,r}$ varies to some degree as it is also a function of porosity (Iglauer et al. 2011b) and pore body-pore throat aspect ratio (Wardlaw et al. 1988, Pentland et al. 2012); and CO₂ was located mainly in larger pores, although in some cases it also spread into relatively small pores. This indicates weakly water-wet conditions, consistent with 2D micromodel (Chalbaud et al 2010, Kim et al. 2012) and contact angle (Chiquet et al. 2007, Broseta et al. 2012, Farokhpoor et al. 2013, Saraji et al 2013, 2014, Kaveh et al 2014, Al-Yaseri et al. 2015, 2015a, Iglauer et al. 2015a, 2015b, Sarmadivaleh et al. 2015) measurements.



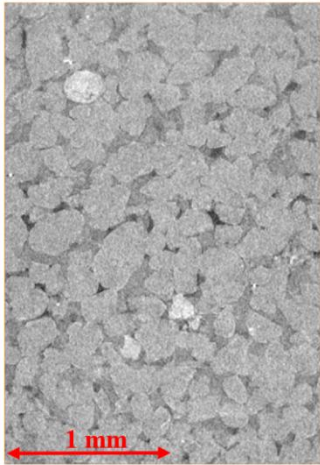
(a)



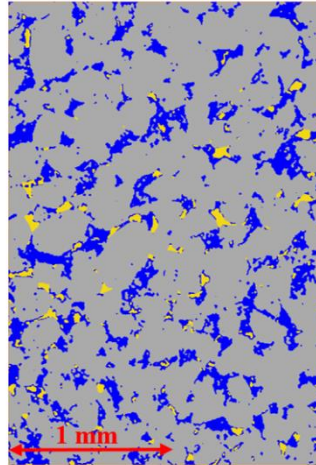
(b)



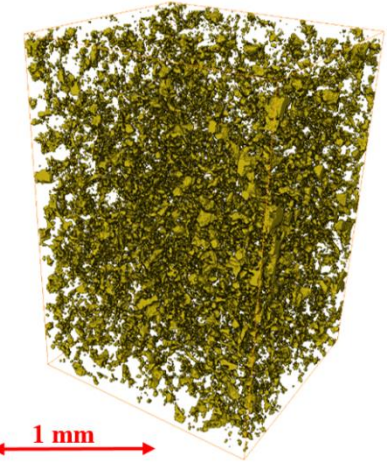
(c)



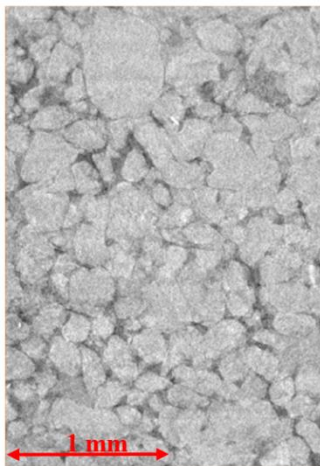
(d)



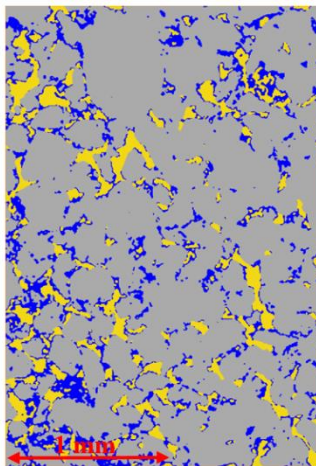
(e)



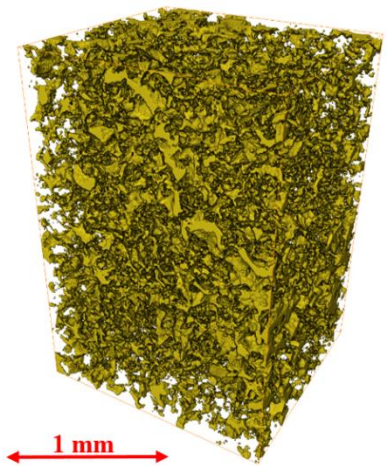
(f)



(g)



(h)



(i)

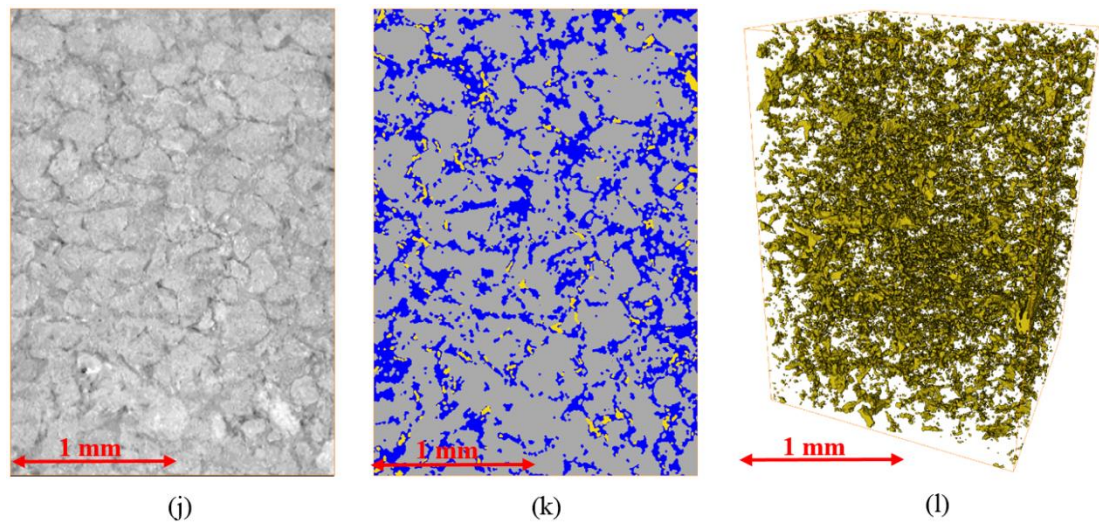


Figure 5-1 μ CT image of Bentheimer sandstone at 10MPa pore pressure and 318K: (a) water-wet initial CO₂ saturation, raw image; (b) water-wet initial CO₂ saturation, segmented image; (c) CO₂ clusters in 3D for the water-wet initial CO₂ saturation, a volume of 3mm³ is shown; (d) water-wet residual CO₂ saturation, raw image; (e) water-wet residual CO₂ saturation, segmented image; (f) CO₂ clusters in 3D for the water-wet residual CO₂ saturation, a volume of 3mm³ is shown; (g) oil-wet initial CO₂ saturation, raw image; (h) oil-wet initial CO₂ saturation, segmented image; (i) CO₂ clusters in 3D for the oil-wet initial CO₂ saturation, a volume of 3mm³ is shown; (j) oil-wet residual CO₂ saturation, raw image; (k) oil-wet residual CO₂ saturation, segmented image; (l) CO₂ clusters in 3D for the oil-wet residual CO₂ saturation, a volume of 3mm³ is shown. CO₂ is black/dark grey, brine is grey and sandstone is light grey; in the segmented images CO₂ is yellow, brine is blue and rock is grey.

Table 5-1 Initial and residual CO₂ saturations in water-wet and oil-wet Bentheimer sandstone at 318K and 10MPa pore pressure.

Wettability state	Contact angle [°]^a	Initial CO₂ saturation [%]	Residual CO₂ saturation [%]
Water-wet	0	50.6	14.9
Oil-wet	130	49.2	8.7

^awater contact angle in air at ambient conditions.

The fact that $S_{CO_2,r}$ was significantly lower in the oil-wet plug (8.7%, i.e. ~40% less CO_2 was trapped in this scenario) is consistent with measurements in an oil-wet plastic bead pack [40], where a substantial reduction of residual $scCO_2$ was observed when compared to a water-wet analogue. Furthermore, CO_2 was frequently located in smaller pores and adjacent to rock surfaces, which indicates intermediate- or CO_2 -wet behaviour. Such wetting behaviour was also observed in 2D micromodels (Chalbaud et al. 2010) and contact angle measurements (Dickson et al. 2006, Yang et al. 2007, 2008, Espinoza et al. 2010, Ameri et al. 2013).

The morphology of the CO_2 clusters in the water-wet plug was also different than in the oil-wet plug (Figure 5-2): while the CO_2 clusters in case of the initial CO_2 saturation state in the water-wet plug were bulky and essentially mimicked the pore space geometry, the corresponding residual CO_2 clusters thinned out considerably and had a somewhat fractal-like shape. Compare this with observations for an oil-water system in a completely water-wet sandstone ((Iglauer et al. 2010); which is known to be strongly water-wet, also at reservoir conditions (Xie et al. 1998)); here the initial and residual clusters mimicked the pore geometry. We conclude that the more fractal-like shape for the residual CO_2 clusters is caused by a shift towards weakly water-wet conditions (CO_2 -brine-quartz is weakly water-wet at reservoir conditions, (e.g. (Al-Yaseri et al. 2015))). The initial CO_2 clusters in the oil-wet plug were already thinner, more “skeleton-like”, than their water-wet counterparts, and the associated residual CO_2 clusters were flat and sheet-like (Figure 5-2d). This is somewhat similar to what has been observed for an oil-water system in an intermediate-wet sandstone (Iglauer et al. 2013), where sheet-like residual oil clusters were observed. This indicates that the oil-wet plug showed intermediate-wet or CO_2 -wet behaviour, consistent with 2D micromodel and contact angle measurements, see above.

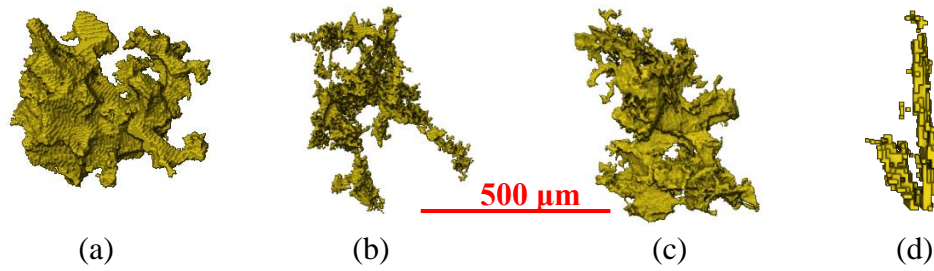


Figure 5-2 Morphologies and volumes of the largest initial and residual CO₂ clusters in Bentheimer sandstone at 318K and 10MPa pore pressure: (a) water-wet initial CO₂ saturation (volume = 27.7 x 10⁶ μm³), (b) water-wet residual CO₂ saturation (volume = 5.39 x 10⁶ μm³), (c) oil-wet initial CO₂ saturation (volume = 18.4 x 10⁶ μm³), (d) oil-wet residual CO₂ saturation (volume = 3.03 x 10⁶ μm³).

5.3.2 CO₂ cluster statistics

5.3.2.1 CO₂ cluster size distributions

CO₂ cluster size distributions were analysed for all scenarios (oil-wet, water-wet, residual, initial), they are plotted in Figure 5-3. As in other experiments of this type (Iglauer et al. 2011a, Iglauer et al. 2010, 2012, 2013, Andrew et al. 2013, 2014, Georgiadis et al. 2013, Geistlinger et al. 2014, 2015) power-law correlations were observed ($N \propto V^{-\tau}$, where N is the count number of clusters of size V , and τ is a fitting exponent), Table 5-2.

Table 5-2 Statistical parameters associated with the initial and residual CO₂ clusters for oil-wet and water-wet sandstone (Bentheimer, 318K, 10MPa pore pressure).

Plug state		τ	\mathbf{p}
Water-wet	Initial CO ₂	1	0.79
	Residual CO ₂	1.6	0.85
Oil-wet	Initial CO ₂	1	0.83
	Residual CO ₂	1.9	0.85

This essentially means that, for each scenario, a large number of small clusters was observed, and only few large clusters. As expected the largest CO₂ clusters at initial CO₂ saturation state ($V_{max} = 28 \times 10^{-6} \mu\text{m}^3$ for the water-wet rock and $V_{max} = 18 \times 10^{-6} \mu\text{m}^3$ for the oil-wet rock) were significantly reduced by waterflooding ($V_{max} = 5 \times 10^{-6} \mu\text{m}^3$ for the water-wet rock and $V_{max} = 3 \times 10^{-6} \mu\text{m}^3$ for the oil-wet rock). Furthermore, the largest clusters always spanned almost the entire observed volume, which may imply that waterflooding (especially in oil-wet reservoirs) reduces the efficiency of structural trapping as the height of the CO₂ plume may not be affected while the CO₂ saturation in the bulk rock is reduced.

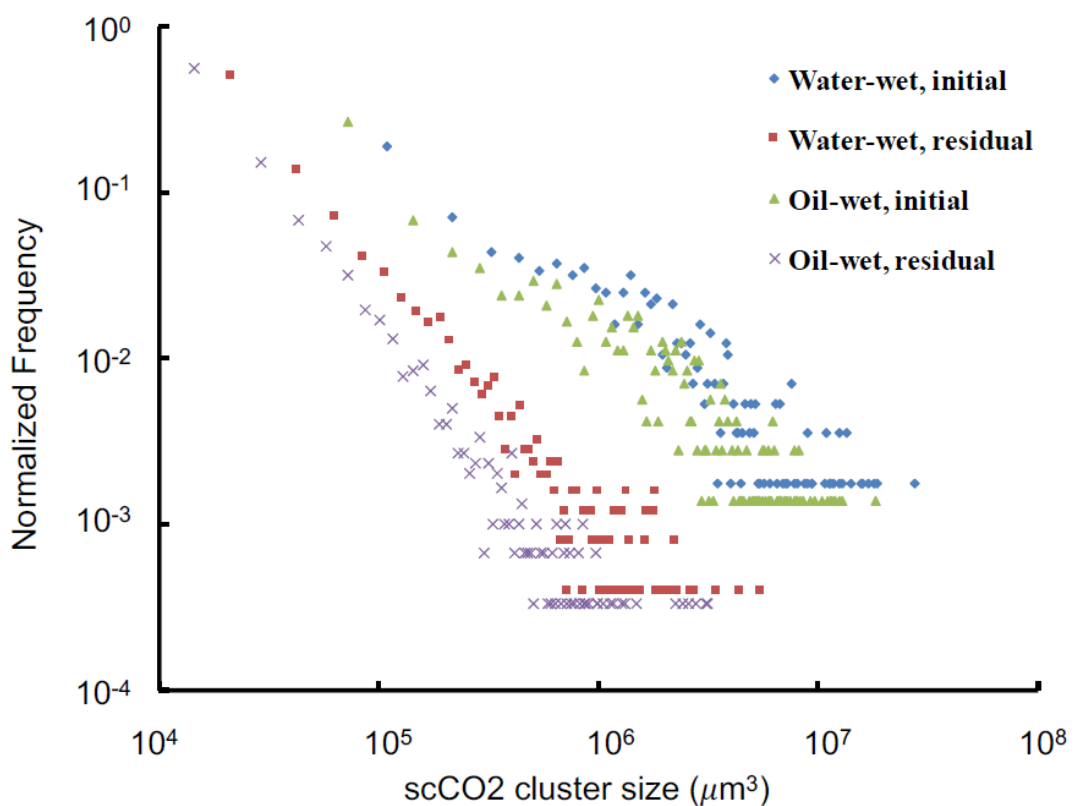


Figure 5-3 Initial and residual CO₂ droplet size distributions measured in water-wet and oil-wet Bentheimer sandstone at 318K and 10MPa pore pressure.

With respect to the scaling exponent; τ was significantly increased by waterflooding; this effect is caused by snap-off of many small CO₂ droplets (Roof 1970). Interestingly, τ increased more in the oil-wet plug, an effect which must be caused by the wettability (as all other parameters were constant, see methods section

above). Mechanistically, CO₂ can enter small pores in the oil-wet plug (due to the lower capillary entry pressure constraint (Oren et al. 1992); as oil-wet surfaces are more CO₂-wet, cp. review paper by (Iglauer et al. 2015a)), and these CO₂ volumes in the smaller pore channels may also snap-off, leading to additional small clusters (= larger τ).

5.3.2.2 CO₂-water interfacial areas

Subsequently, surface areas A for each CO₂ cluster were measured and plotted against their volumes V , see Figure 5-4. From the graph, again a power-law relation is evident ($A \sim V^{-p}$); and as with all other measurements of this kind (Karpyn et al. 2010, Pentland et al. 2012, Iglauer et al. 2013), p was always ~ 0.8 .

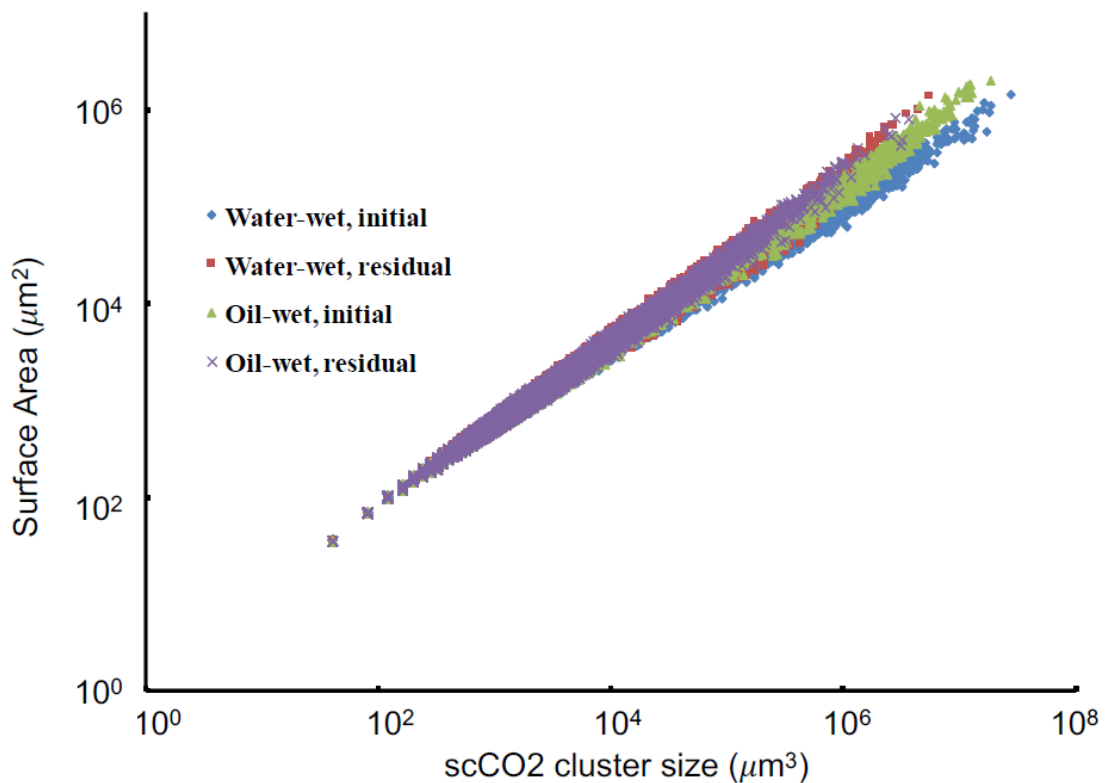


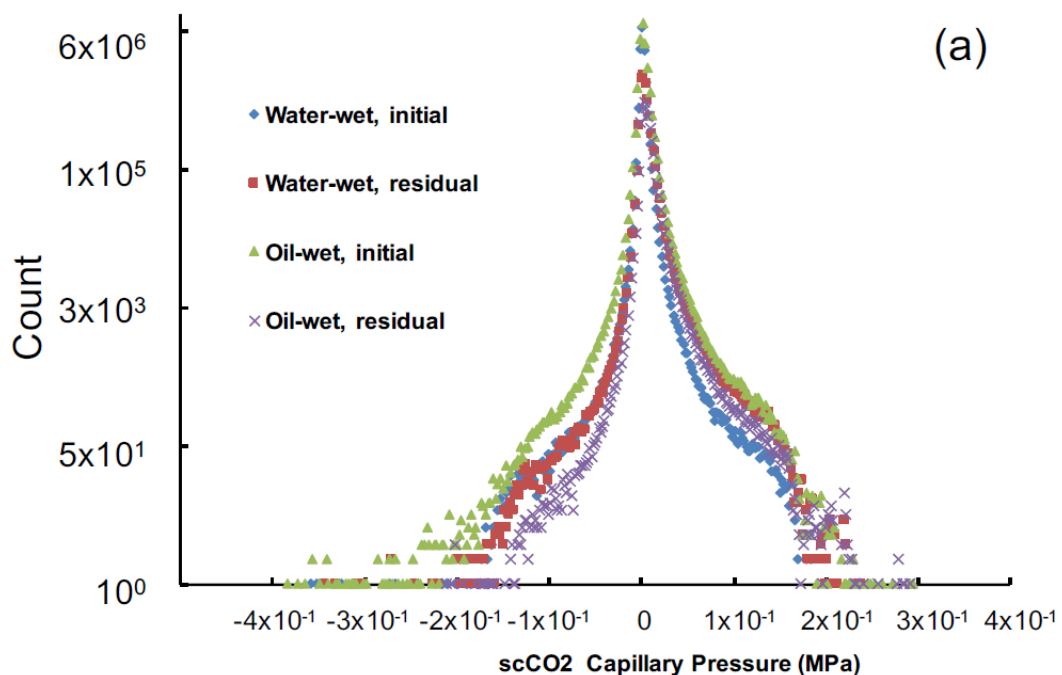
Figure 5-4 CO₂ droplet surface areas plotted against their volumes for various conditions, at 318K and 10MPa pore pressure.

In this context, it is interesting to note that the power-law exponents measured (τ and p) are significantly lower than predicted by percolation theory (which predicts $\tau = 2.189$ and $p \approx 1$; (Stauffer 1979, Lorenz 1998)). We conclude that simple percolation models will not deliver accurate and reliable predictions, as already evident in Georgiadis et al.'s work (Georgiadis et al. 2013).

5.3.2.3 Capillary pressure distributions

Finally, we measured curvatures (C) of each cluster, and calculated associated capillary pressures (p_c) for each CO₂ bubble and water droplet ($p_c = \gamma C$; Figure 5-5) assuming a water-CO₂ interfacial tension γ of 40mN/m (Li et al. 2012, Arif et al. 2016).

The capillary pressure distributions were roughly bell-tower shaped, and always showed a narrow and high peak around zero capillary pressure, consistent with data reported for oil-water systems (Armstrong et al. 2012). Nominally the capillary pressures ranged from ~ -0.4 MPa to $+0.3$ MPa in case of CO₂, and from -0.6 MPa to $+0.6$ MPa in case of water (the water distributions were slightly stretched to the left, Figure 5-5b).



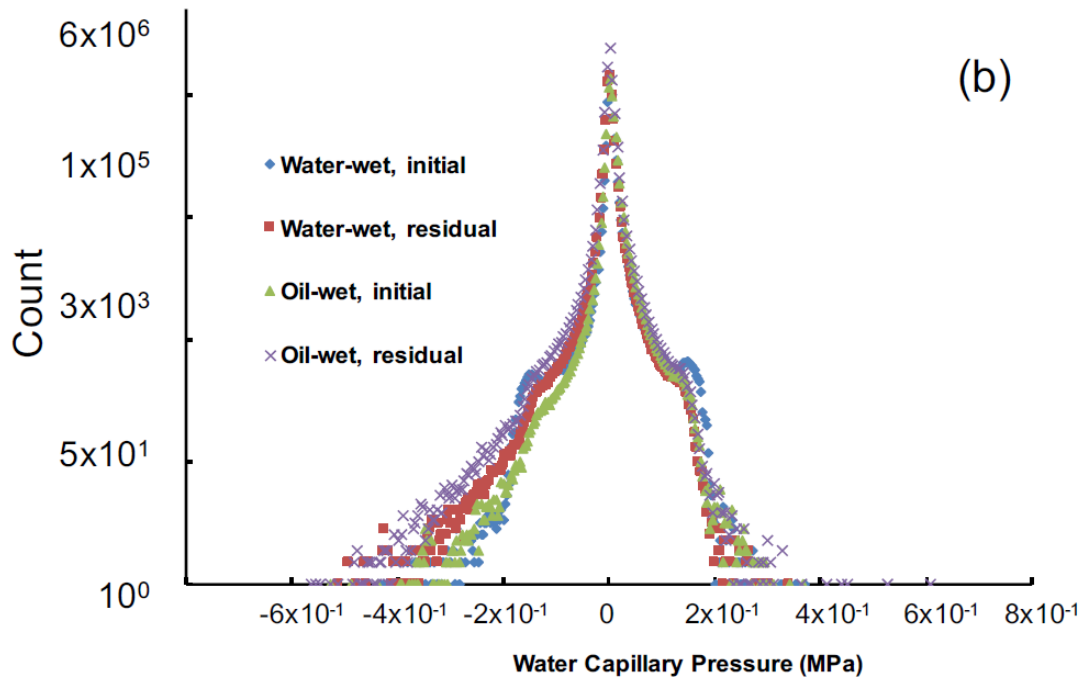


Figure 5-5 CO₂ (a) and water (b) droplet capillary pressures measured in water-wet and oil-wet Bentheimer sandstone at 318K and 10MPa pore pressure.

5.4 Conclusions and Implications

We demonstrated that residual CO₂ saturations ($S_{CO_2,r}$) in oil-wet sandstone are dramatically lower than in equivalent water-wet sandstone ($S_{CO_2,r} = 8.7\%$ versus 14.9%). This is consistent with data reported for a plastic bead pack (Chaudhary et al. 2013), and indicates that oil-wet rock is intermediate-wet or CO₂-wet at reservoir conditions; a phenomenon which was also observed in independent contact angle measurements (Dickson et al. 2006, Yang et al. 2007, 2008, Espinoza et al. 2010, Ameri et al. 2013). This shift in CO₂-wettability also significantly influenced CO₂ cluster morphologies and residual cluster size distributions: flatter and more sheet-like clusters and more small CO₂ bubbles were observed in the oil-wet sample.

We conclude that, importantly, residual trapping is less efficient in oil reservoirs than in deep saline aquifers.

Chapter 6 Influence of wettability on residual gas trapping and enhanced oil recovery in three-phase flow: a pore-scale analysis using micro-computed tomography*

Abstract

We imaged an intermediate-wet sandstone in 3D at high resolution ($1\text{-}3.4\mu\text{m}$)³ with an x-ray micro-computed tomograph at various saturation states. Initially the core was at connate water saturation and contained a large amount of oil (94%), which was produced by a waterflood (recovery factor $R_f = 52\%$ OOIP) or a direct gas flood ($R_f = 66\%$ OOIP). Subsequent water and/or gas flooding (WAG process) resulted in significant incremental oil recovery ($R_f = 71\%$ OOIP), while a substantial amount of gas could be stored ($\sim 50\%$) – significantly more than in an analogue water-wet plug. The oil and gas cluster size distributions were measured and followed power-law correlations $N \propto V^{-\tau}$ (where N is the frequency with which clusters of volume V are counted) with decay exponents τ between 0.7 and 1.7. Furthermore, the cluster volume V plotted against cluster surface area A also correlated with a power-law $A \propto V^p$, and p was always ≈ 0.75 . The measured τ and p values are significantly smaller than predicted by percolation theory (which predicts $p \approx 1$ and $\tau = 2.189$); which raises increasing doubts about the applicability of simple percolation models. In addition, we measured curvatures and capillary pressures of the oil and gas bubbles in-situ, and analysed the detailed pore scale fluid configurations. The complex variations in fluid curvatures, capillary pressures and the fluid-fluid or fluid-fluid-fluid pore scale configurations (exact spatial locations also in relation to each other and the rock surface) are the origin of the well-known complexity of three-phase flow through rock.

*Reference: Iglauer et al. (2016) in Journal of Society of Petroleum Engineers, SPE 179727

6.1 Introduction

Three-phase flow through porous media is an important topic in the oil and gas industry as it occurs in gas-assisted enhanced oil recovery (EOR) processes, and carbon capture and storage (CCS), when CO₂ is injected into oil reservoirs for storage and incremental oil recovery (Blunt et al. 1993, IPCC 2005); furthermore, three-phase flow also occurs in soils contaminated by organic solvents (Sleep and McClure 2001). In this context, a significant amount of work has been completed to study such three-phase flow at the reservoir scale (e.g. compare Qi et al. 2009, Moortgat et al. 2011, Beygi et al. 2013) and the mesoscale (centimetre scale; e.g. Oak et al. 1990, Egermann et al. 2000, Caubit et al. 2004, Al-Mansoori et al. 2009, Amaechi et al. 2014). However, the detailed pore-scale (micrometre scale) fluid dynamics are poorly understood despite the fact that pore-scale flow determines the overall reservoir-scale (kilometre scale) flow, in addition to other larger scale mechanisms, e.g. reservoir heterogeneity or gravity segregation (Bear, 1988). It is therefore vital to deepen our understanding in this area to aid the development of better CCS and EOR schemes and to improve clean-up technologies for contaminated soils.

Such pore-scale three-phase flow has previously been studied theoretically (e.g. Van Dijke et al. 2011; Piri and Blunt 2004, 2005a,b; Al-Dhahli et al. 2012) or experimentally using 2D micromodels (e.g. Oren 1992, Oren and Pinczewski 1995, Dong et al. 1995, Keller et al. 1997); both approaches rely on strongly simplified pore morphologies. It is, however, well established that the pore space morphology has a significant impact on the related fluid dynamics (Dullien 1979); in addition, it is clear that 3D flow characteristics cannot be fully captured by 2D models (e.g. the percolation threshold is significantly lower in 3D, Stauffer 1979). Only recently, with the advent of micro-computed tomography (μ CT), three-phase flow through porous media can be studied in 3D at the pore-scale of realistic geological porous media (compare the recent reviews by Blunt et al. (2013), Wildenschild and Sheppard (2013) and Cnudde and Boone (2013)), which is most relevant for industrial applications.

In this work, we extend our previous investigation of three-phase pore-scale flow in water-wet sandstone (Iglauer et al. 2013) to include intermediate-wettability. This wettability effect is of particular importance in sandstone and carbonate oil reservoirs as these are frequently intermediate-wet or even oil-wet (Cuiec 1991). Specifically, wettability has a strong impact on petrophysical parameters (Anderson 1987) including residual saturations (Morrow 1990, Tanino and Blunt 2013), relative permeabilities (McCaffery and Bennion 1974), capillary pressures (Morrow 1976, Anderson 1987), fluid topologies (Iglauer et al. 2012), and CO₂ storage capacities (Iglauer et al. 2015). We demonstrate that there is a large variation in fluid configurations, fluid-fluid curvatures and capillary pressures, which leads to the well-known complexity of three-phase flow (e.g. cp, Oren et al. 1992).

6.2 Experimental Methodology

We selected a clean homogenous water-wet outcrop sandstone (Clashach, from Elgin in Scotland), and from a larger block we drilled a small cylindrical plug (5mm diameter, 9mm length). The sandstone had a porosity of 10%, Table 6-2, and a brine permeability of $9.6 \times 10^{-14} \text{ m}^2$ (Iglauer et al. 2012); its chemical composition was measured via XRD: 96wt% quartz, 2wt% K-feldspar, 1wt% calcite, 1wt% ankerite and illite traces. The plug was mounted in a flow cell, and its wettability was altered with North Sea crude oil using a procedure described previously (Iglauer et al. 2012) to intermediate-wet (Amott-Harvey index = -0.1, Amott oil index = -0.1, and Amott water index = 0; which is slightly oil-wet). The aged plug (now mounted in a flow cell) was then imaged with a micro-computed tomograph (Xradia Versa-XRM-500) at a resolution of $(3.4 \mu\text{m})^3$. Note that an inner cylindrical volume of the plug was imaged with a diameter of 3.4mm to achieve good resolution (i.e. not the whole 5mm-diameter plug). The same resolution was used throughout the whole work flow for each μCT scan except for a few high-resolution scans ($\sim 1 \mu\text{m}$)³ with which we investigated the detailed fluid configurations in the pore space. The aged plug was vacuumed for 60 min (to ensure no air was trapped in the pore space) and deaired brine was injected for 5 min under vacuum followed by injection of ~ 1000 pore volumes (PV) of oil at a capillary number ($= \mu q / \gamma$, where q is the Darcy flow rate, μ is the viscosity of the injected phase and γ is the interfacial tension) of 3.54×10^{-4} to

achieve connate water saturation S_{wc} . The brine was a 10wt% solution of potassium iodide (KI) in deionized water, the KI is required to achieve sufficient x-ray contrast between the fluids. The oil used was 1-bromododecane (purity $\geq 99\text{mol}\%$, from Sigma Aldrich), again to enable CT contrast. The plug then essentially represented an oil reservoir (“connate water saturation”, with no gas present), and we acquired a μCT image of the plug in this state. Following our previous study line (Iglauer et al. 2013), we then distinguished two cases to study residual gas and oil topologies and distributions: the “waterflood” scheme resembles the process typically used in the petroleum industry, where oil reservoirs are waterflooded to enhance oil recovery (Lake 1990), and subsequently gas flooded for further improving oil recovery (e.g. Blunt et al. 1993) and/or gas storage (e.g. CO_2 storage, cp. e.g. Iglauer et al. 2011a, Pentland et al. 2012, Iglauer et al. 2013), i.e. a WAG flood. The second flow scheme “gasflood” uses an alternative process sequence, which has been shown to be significantly more efficient in terms of oil recovery and gas storage in a water-wet sandstone (Iglauer et al. 2013).

6.2.1 Waterflood

25 PV of brine were injected at a capillary number of 7.4×10^{-7} into the plug at connate water saturation to establish residual oil saturation (S_{or}); subsequently, approximately 50 PV of N_2 gas were injected at S_{or} at a pressure drop of 20000 Pa over a 9mm plug length (0.05 mL/min flow rate for 6 minutes), which approximately corresponded to a capillary number of 0.01, so that an initial gas saturation state ($S_{gi,bg}$) was established. This was followed by another brine flood (using the same capillary number as above, $\sim 7.4 \times 10^{-7}$) to measure residual oil and gas saturations and configurations ($S_{gr,bgb}$ state).

6.2.2 Gasflood

N_2 gas was directly injected into the plug at connate water saturation (S_{wc}) using the same conditions as above to establish the initial gas saturation, $S_{gi,g}$. This was followed by brine injection (same conditions as above) to establish residual gas saturation, $S_{gr,gb}$.

All plug states were μ CT imaged; and all experiments were conducted at ambient laboratory conditions (295K and 0.1MPa). Note that the plugs were always positioned horizontally during the floods. Bond numbers were estimated to be $\approx 10^{-4}$ for the liquid-gas and the oil-brine system; we therefore do not expect the residual clusters/residual saturations to be influenced by buoyancy forces (Morrow et al. 1988).

We note that the oil used was non-spreading with a spreading coefficient S ($= \gamma_{gw} - \gamma_{ow} - \gamma_{go}$, where γ_{gw} is the gas-brine interfacial tension, γ_{ow} is the oil-brine interfacial tension and γ_{go} is the gas-oil interfacial tension, cp. Table 6-1) of -11.05 mN/m; and it is clear that the algebraic sign of S has a profound influence on the fluid dynamics and statics (Blunt et al. 1995, Oren and Pinczewski 1995).

Table 6-1 Interfacial tensions of the fluids used.

Fluid-fluid system	Interfacial tension [mN/m]
water/1-bromododecane *	$\gamma_{ow} = 52.09$
water/nitrogen **	$\gamma_{gw} = 72$
1-bromododecane/nitrogen ***	$\gamma_{go} = 30.96$

* measured at 295K (Ortiz-Arango and Kantzas 2009).

** measured at 293.15K and 0.101 MPa (Lide 2007).

*** surface tension of 1-bromododecane.

6.3 Results and Discussion

Figure 6-1 shows a set of 2D slices through the rock and pore space filled with the various fluids at different saturation stages: connate water saturation (S_{wc}), residual oil saturation (S_{or}), the two initial gas saturations ($S_{gi,g}$ and $S_{gi,bg}$), and the two residual gas saturations ($S_{gr,gb}$ and $S_{gr,bgb}$). The slices display a greyscale map, where each greyscale corresponds to a specific relative radiodensity of a volume element, and this greyscale can be assigned to a certain material (i.e. rock, oil, gas or water),

Golab et al. (2013). The gas has the lowest relative radiodensity and it is thus black, while oil (because it is brominated) has the highest relative radiodensity: it is white. The brine and the solid grains have a relative radiodensity in between, they are both grey; the brine being slightly darker than the grains. On the left side greyscale image slices were shown, which were filtered with a non-local means filter, Buades et al. (2005), and segmented with a watershed algorithm (Schlüter et al. 2014; right side). The detailed pore-scale fluid configurations can be observed at this resolution ($\sim 1\text{-}3\ \mu\text{m}$), this is discussed further below; moreover, the 3D topology of the oil and gas phases can be visualized (Figure 6-2) and analysed.

Porosities, and oil, water and gas saturations were subsequently measured on the segmented μCT images (Table 6-2). The oil recovery factor R_f is also shown in Table 6-2, it has been calculated based on the initial oil in place ($R_f = (S_{o,wc} - S_o) / S_{o,wc}$ with S_o = actual oil saturation, $S_{o,wc}$ = oil saturation at S_{wc}). Initially the oil saturation at S_{wc} was quite high, because a high capillary pressure was applied; consequently, a relatively high residual oil saturation was obtained after waterflooding as a high initial oil saturation leads to a high residual saturation (e.g. compare Pentland et al. 2010, 2011, Iglauer et al. 2011b, Gittins et al. 2010).

Gas injection directly into the plug at connate water saturation or into the plug at residual oil saturation both produced significant amounts of incremental oil, consistent with literature results (e.g. Dumoré and Schols 1974, Egermann et al. 2000, Al-Mansoori et al. 2009, Amaechi et al. 2014). The final waterflood again produced a significant amount of additional oil. In the intermediate-wet plug, oil recovery efficiencies for both flood sequences were similar (“waterflood”: $R_f = 70.8\%$ OOIP; “gasflood”: $R_f = 71.1\%$ OOIP). Compare this with recovery factors in an analogue water-wet plug (“waterflood”: $R_f = 68.9\%$ OOIP; “gasflood”: $R_f = 77.1\%$ OOIP), here the gas flood was significantly more efficient, cp. Table 6-2.

We furthermore observed that the residual gas saturation was similar in both flooding processes in the intermediate-wet plug, while in the analogue water-wet plug, direct gas flooding led to higher residual gas saturations, Table 6-2. In the context of CO_2 storage, it is of key relevance to estimate how much CO_2 can be stored: for residual trapping, this can be accomplished with the capillary trapping capacity C_{trap} , which

quantifies how much gas can be stored in a unit rock bulk volume (Iglauer et al. 2011b), C_{trap} is defined as the product of rock porosity and residual gas saturation ($= \phi S_{\text{gr}}$). Interestingly C_{trap} is much higher for the intermediate-wet plug; this is contrary to what has been observed in two-phase systems, where residual gas saturation is lower in intermediate-wet systems (cp. Iglauer et al. 2011a with Chaudhary et al. 2013).

Table 6-2 Porosities, oil S_o , water S_w , and gas saturations S_g measured on the μ CT images (percentages shown). * water-wet condition (taken from Iglauer et al. 2013). R_f is the oil recovery factor, C_{trap} is the capillary trapping capacity.

Saturation state	Porosity	S_o	S_g	S_w	C_{trap}	R_f [OOIP]
S_{wc}	9.3	94.3	0	5.7	---	0
S_{or}	10.4	45.3	0	27.8	---	52.0
$S_{\text{gi,g}}$	11.6	32.2	48.8	19.0	4.88	65.8
$S_{\text{gi,bg}}$	9.8	35.0	56.8	8.2	5.68	62.9
$S_{\text{gr,gb}}$	10.1	27.2	54.6	18.2	5.46	71.2
$S_{\text{gr,bgb}}$	10.5	27.5	50.0	22.5	5.00	70.8
$S_{\text{gr,gb}}^*$	11.6	21.6	20.6	57.8	2.06	77.1
$S_{\text{gr,bgb}}^*$	10.7	29.3	5.3	65.4	0.53	68.9

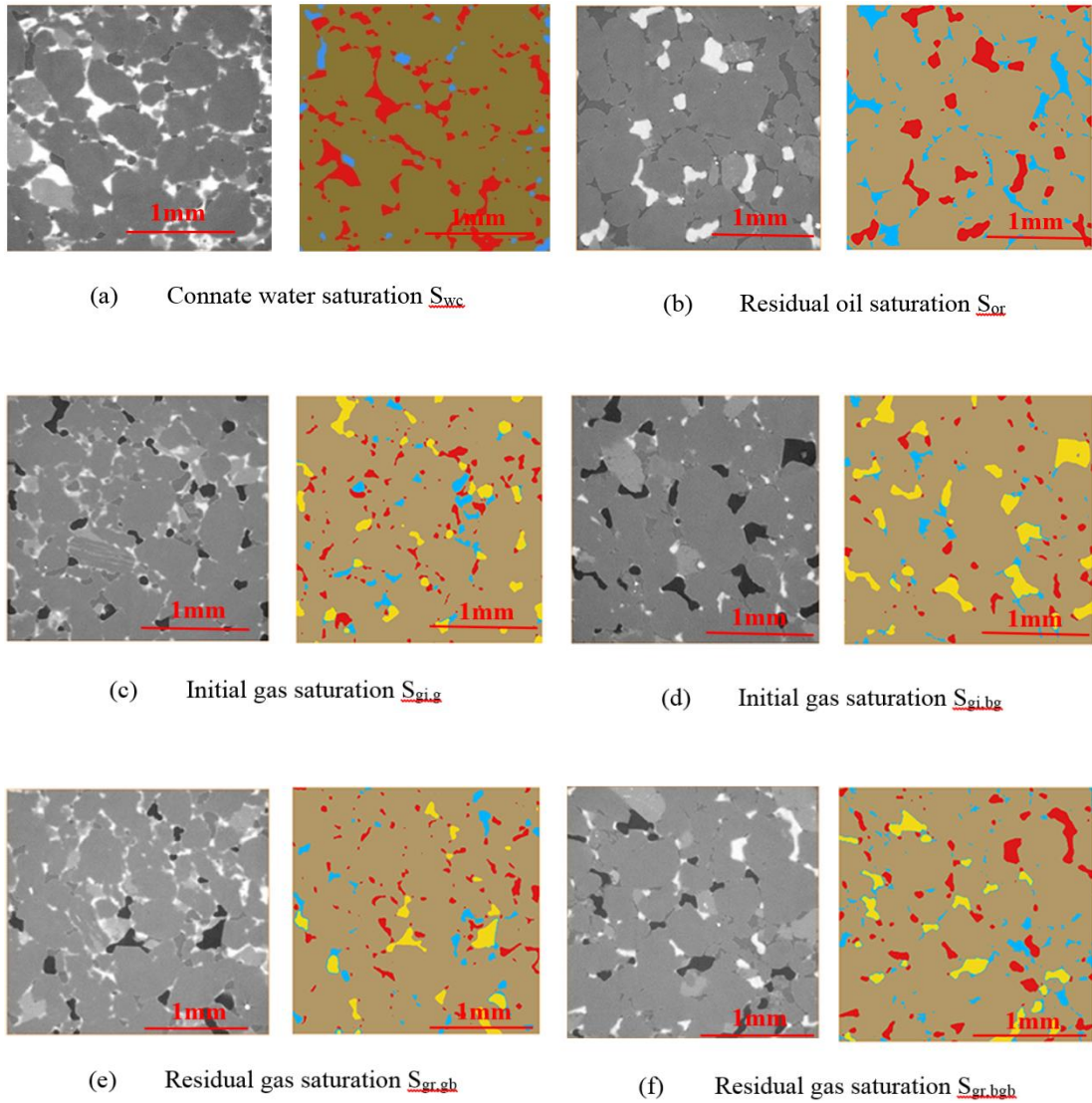


Figure 6-1 2D slices through the rock and various fluids at different saturation states. (a) connate water saturation S_{wc} ; (b) residual oil saturation S_{or} ; (c) initial gas saturation $S_{gi,g}$; (d) initial gas saturation $S_{gi,bg}$; (e) residual gas saturation $S_{gr,gb}$; (f) residual gas saturation $S_{gr,bgb}$. Left: greyscale images; right: segmented images. In the raw images, oil is white, gas is black, brine dark grey, and sandstone is light grey. Oil is red, brine blue, gas yellow and rock is brown in the segmented images.

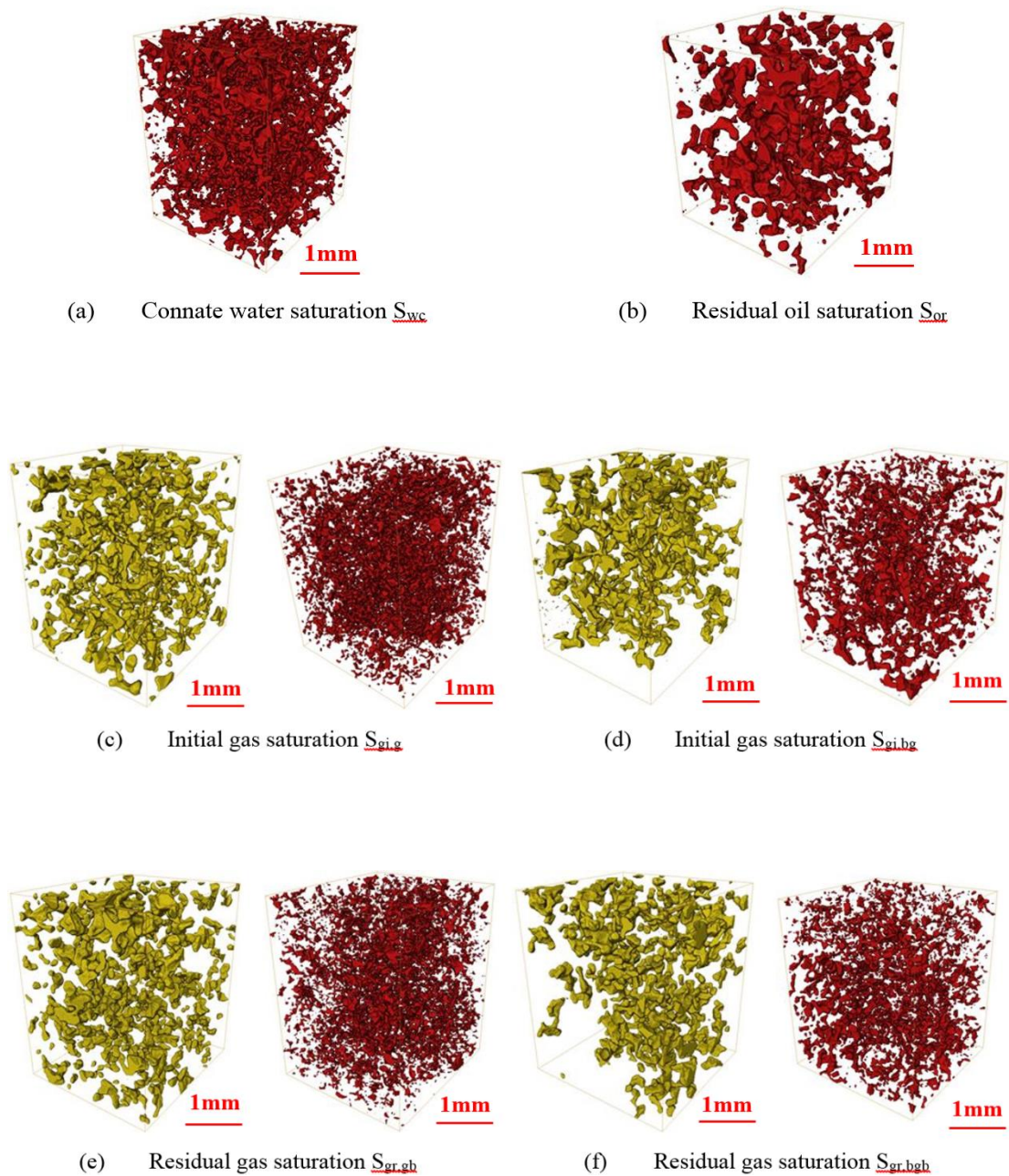


Figure 6-2 3D topology of the gas (yellow), oil (red) phases at various saturation states. All volumes shown are $\sim 12\text{mm}^3$.

Figure 6-2 clearly demonstrates that the oil saturation decreases substantially during the initial waterflood (from the S_{wc} to the S_{or}), and S_o is further reduced in the tertiary production steps. The oil is uniformly distributed but split into many separate clusters; see below for a more detailed analysis. Moreover, it is apparent from the images that the gas saturation S_g did not change significantly, and the gas was also distributed in many individual clusters.

6.3.1 Qualitative analysis Pore-scale fluid configurations

The wettability state of the sandstone influenced oil distribution and fluid topology in the pores as expected and measured previously in two-phase oil-water systems (Kumar et al. 2009, 2012, Iglauer et al. 2012).

6.3.1.1 Waterflood

The oil saturation was high at S_{wc} , because of the high applied capillary pressure during primary drainage. However, many small and some large water volume elements were still present in the pore network of the rock; the curvature of these ganglia varied quasi randomly from convex to concave throughout the plug, apparently depending on the wettability state of the directly adjacent pore surfaces; this is discussed further below. Moreover, oil and water were both present in the smallest pores, which indicates intermediate wettability as expected and consistent with the independent Amott index measurement (see above).

The waterflood was very efficient and about half of the oil was produced (S_o decreased from 94.3% at S_{wc} to 45.3% at S_{or} , Table 6-2); in addition, a significant amount of oil was displaced from the smaller pores, which again indicated intermediate-wet conditions. At S_{or} , oil was located mainly in medium-sized and larger pores, although some oil still resided in smaller pores. Compare this with a water-wet plug: here all residual oil is located in large pores as disconnected bubbles (Iglauer et al. 2010). We infer that oil was partially mobilized into oil banks during production, consistent with mesoscopic observations (e.g. Egermann et al. 2000, Gittins et al. 2010).

A similar situation was observed at the $S_{gr, bgb}$ state where oil had been displaced from small pores to a large extent (but not fully), and was trapped in medium-sized pores, probably by snap-off events (Roof 1970). We note that oil layer flow (cp. for instance Table 6-3 below) may contribute to further oil recovery, but at a slow rate, because of the low relative permeability of these layers (caused by their thin cross-sectional area) (e.g. Oren et al. 1992, and Blunt et al. 1995). In this context one clear limitation of the μ CT technique needs to be mentioned: thin layers (thickness of a

few to a few hundred nanometres) cannot be observed, although they play a significant role in terms of wettability effects and associated fluid dynamics (e.g. Hirasaki 1991, Blunt et al. 1995, Tokunaga and Wan 2013).

In summary, the oil phase was often touching the rock surface (i.e. oil-wet behaviour), but in many cases oil drops were also located in pore centres, an indication of water-wet behaviour – this adds up to an intermediate-wet state. In contrast, oil is always sandwiched between the gas and water phases in water-wet rock, cp. Figure 6-3 (Iglauer et al. 2013).

6.3.1.2 Gas Injection

More oil accumulated in medium-sized and larger pores in the $S_{gi,bg}$ state than in the $S_{gi,g}$ state (due to the additional waterflooding step), while more oil was present in small pores in the $S_{gi,g}$ state. These configurations somewhat reflect the oil distributions prior to gas flooding. However, in both sequences - more ($S_{gr,gb}$) or much more ($S_{gr,gb}$ state) - oil was present in smaller pores in the intermediate-wet plug than in the analogue water-wet plug, Figure 6-3, where oil was mainly localized in larger pores.

In both flood sequences, however, gas entered the largest pores, and was typically located in their centres (in both, the intermediate-wet and the water-wet plugs; this is expected as gas is the most non-wetting phase in both cases, Oren and Pinzcewski (1995) Consequently, the gas bubbles were frequently larger than the oil clusters, cp. Figures 6-2. In addition, the gas clusters adopted a range of morphologies: from quasi-spherical to ganglia-like; their specific morphologies were determined by pore connectivity, pore geometry, rock wettability and capillary pressure p_c (Youssef et al. 2010, Iglauer et al. 2012, Berg et al. 2013).

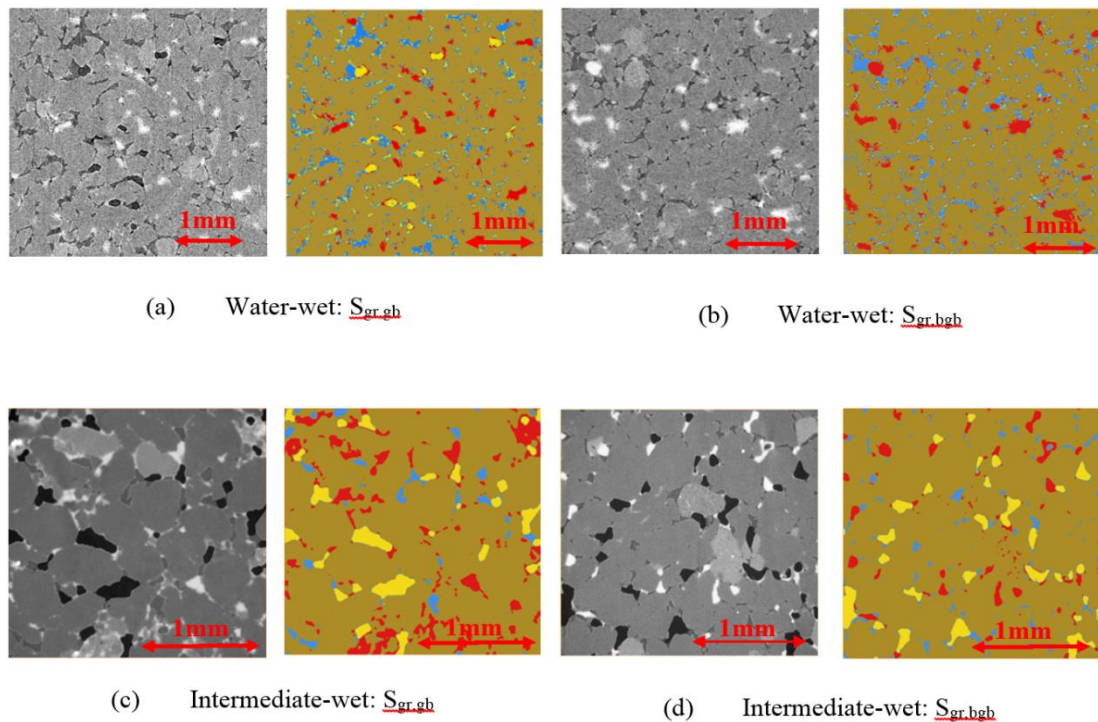


Figure 6-3 Comparison of oil and gas phase configurations at the two residual gas saturation states ($S_{gr,gb}$ and $S_{gr,bgb}$) in water-wet (top) and intermediate-wet (bottom) Clashach sandstone. Note that the plugs were drilled from the same block. The water-wet plug images were acquired with monochromatic light at a nominal resolution of $(9\mu\text{m})^3$ (Iglauer et al. 2013); the intermediate-wet plug images are those discussed in this script, their resolution is $(3.4\mu\text{m})^3$.

Figure 6-4 shows slices through high resolution μCT images for some selected areas for both residual gas saturation states; the pore-scale configuration of the different fluid phases can be quite clearly seen. Most configurations observed were previously identified theoretically (Zhou and Blunt 1998, Piri and Blunt 2004), and we can provide experimental verification for these here; in addition, μCT imaging enables a detailed analysis and new subcategories were found (Table 6-3). Furthermore, analysis of the interplay between the pore morphology and surface-free energies (oil-wet or water-wet) and how that affects the exact fluid configurations can be explored. The images demonstrate that these settings are more complicated than current models, which are based on strongly simplified pore geometries and wettability alteration concepts (typically it is assumed that the surface which contacted crude oil changed towards oil-wet, cp. e.g. Piri and Blunt 2004), however,

surfactant molecules can diffuse through the water and adsorb onto the rock surface, which would then also alter the wettability of the surface without direct contact (Buckley et al. 1998). Interestingly the fluid-fluid curvature can be observed (cp. Table 6-3); such curvatures are rather important as they determine pore-scale fluid dynamics (e.g. Jettestuen et al. 2013). These effects are analysed in more detail in the statistics section below.

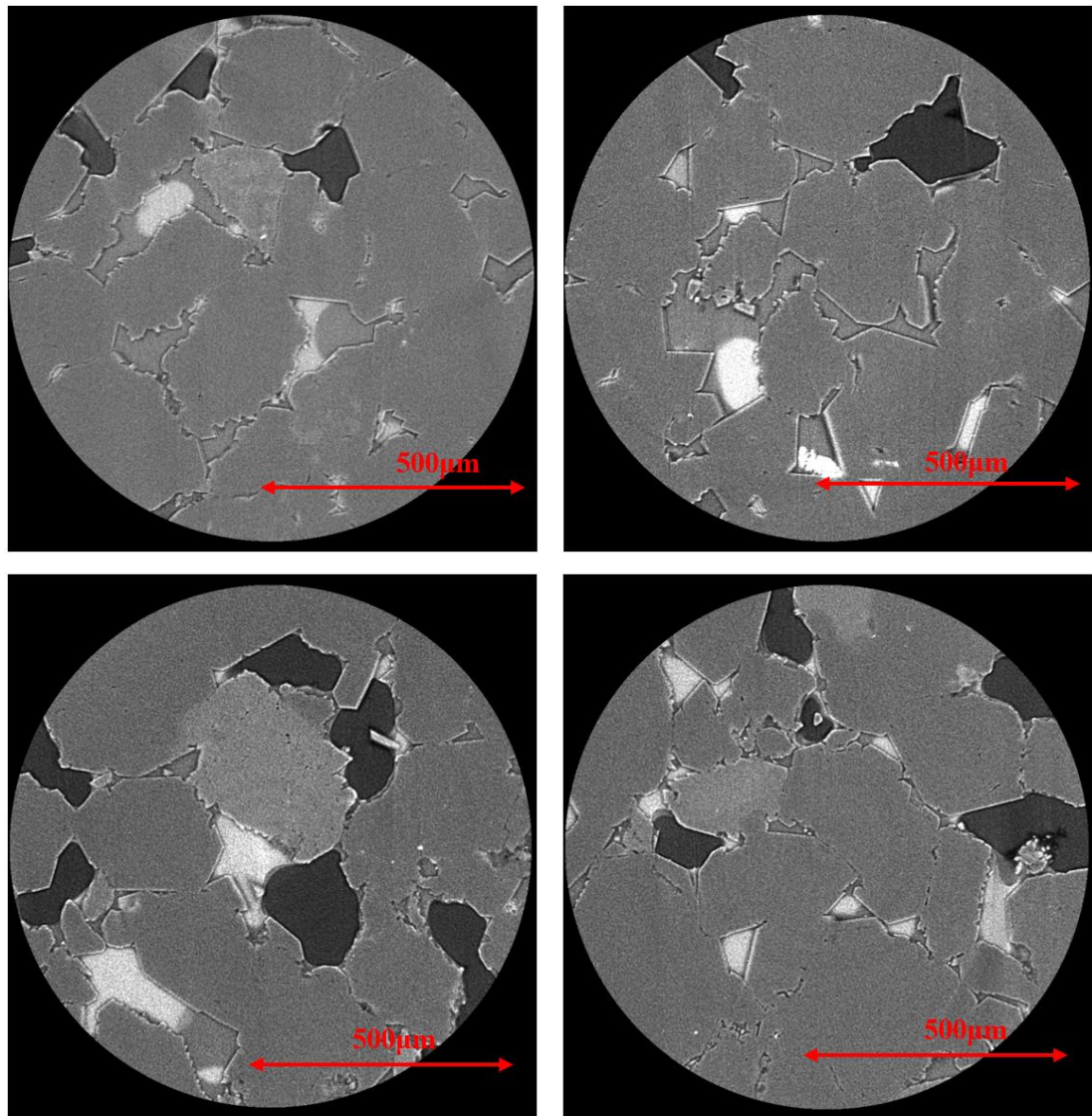



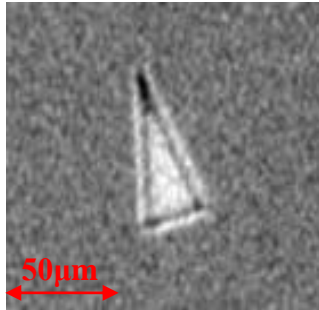

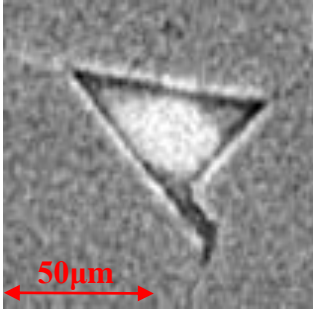

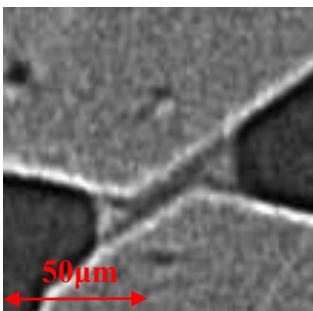

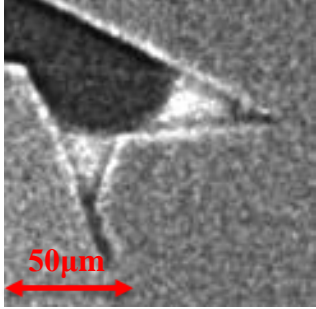

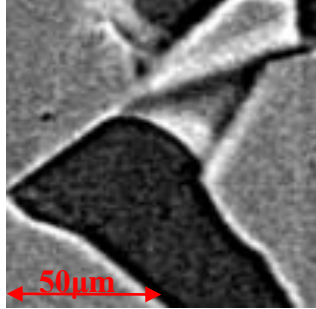

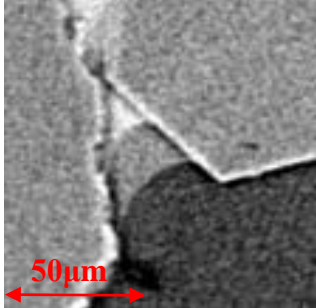

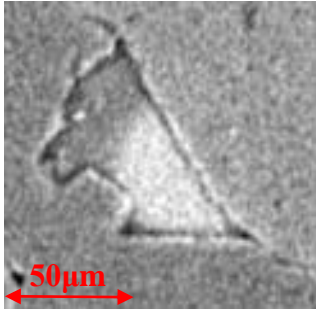

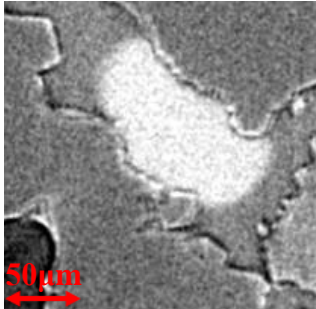

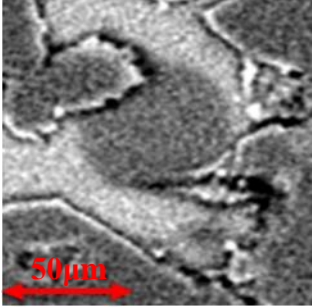

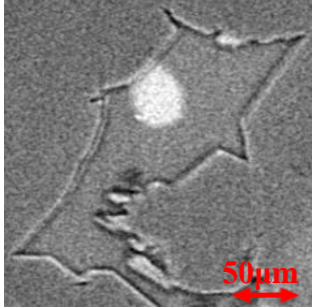

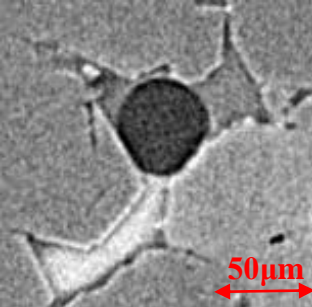


Figure 6-4 Selected high resolution 2D slices through the plug at residual gas saturation; top: $S_{gr,gb}$, $(1.00\mu\text{m})^3$ resolution; bottom: $S_{gr,gb}$, $(1.07\mu\text{m})^3$ resolution. Gas is black, oil white, brine dark grey and rock light grey. For classification consult Table 6-3.

Table 6-3 Pore-corner fluid configurations identified in the μ CT images of the intermediate-wet plug.

Pore configuration			Pore Image at $(1\mu\text{m})^3$ Resolution
Symbol	State	Category	
	Water (1 phase)	W	
	Oil (1 phase)	O	
	Oil in water (2 phase)	WO	
	Gas in water (2 phase)	WG	

	Gas in oil (2 phase)	OG	
	Gas in oil in water (2 phase)	WOG	
	Gas in water in oil (3 phase)	OWG	
	Water and oil wet	f-WO	
	Oil and water: oil curvature = convex	f-WOcx	

	<p>Oil and water: oil curvature = concave</p> <p>f-WOcv</p>	
	<p>Oil droplet attached to rock surface in water</p> <p>f-WOd</p>	
	<p>Water and oil wet with gas in center</p> <p>f-GWO</p>	

The fluid configurations listed in Table 6-3 include pores filled with a single fluid, two fluids or three fluids, and categories for several fractionally-wetted pores. The precise fluid configuration in these fractionally-wetted pores is a consequence of the distribution of oil-wet rock surfaces, for instance gas can be surrounded by oil on one side and brine on another; oil can be located in a pore corner while the remaining pore space is filled with water; or a water drop can stretch from one pore surface to the opposite side (between two isolated water-wet surfaces), surrounded by a concave oil volume, which fills most of the remaining pore space. Furthermore, small water droplets can reside in water-wet locations within a pore.

An interesting phenomenon is that some pores contain convex oil drops which stretch from one oil-wet rock surface to the opposite (oil-wet) side, while others show concave oil drops. This is a clear indication of variation in oil contact angles within the same plug, which again reflects the intermediate-wet character of the

sandstone. Individual oil drops attached to one pore wall side only, and water layers sandwiched between gas and oil in pore corners were also observed: both an indication of partially oil-wet rock walls.

In summary, the large variability in terms of pore-scale fluid configurations adds to the high complexity of three-phase flow (and static distributions) in intermediate-wet rock, and the associated difficulty in predicting such flow with confidence, not only at hectometre reservoir-scale, but already at the micrometre pore-scale. In the next section, these physical phenomena were quantitatively analysed.

6.3.2 Quantitative analysis - Cluster statistics

6.3.2.1 Cluster interfacial areas

The relationships between oil and gas cluster volumes V and their surface areas A are plotted in Figures 6-5 and 6-6. The relations can be very well described using a power-law correlation $V \propto A^p$, where p is around ~ 0.75 for all saturation states, Table 6-4. This is essentially the same value as measured for oil and gas clusters in a water-wet sandstone (Iglauer et al. 2013), and in two-phase flow for residual oil or supercritical CO_2 clusters (Iglauer et al. 2010, 2011, Pentland et al, 2012). We thus hypothesize that $p \sim 0.75$ may be a universal constant, at least for relatively homogeneous sandstones and water-wet to intermediate-wet conditions.

We also note that p is significantly lower than the value predicted by percolation theory (which predicts $p \approx 1$, Stauffer 1979), and this has important implications for mass transfer processes as the real oil-water or gas-water interfacial area is significantly smaller than expected, and associated kinetics (e.g. CO_2 dissolution trapping as a method to dispose CO_2 , Hirai et al. 1997, Suekane et al. 2006, Iglauer 2011), are much slower than anticipated.

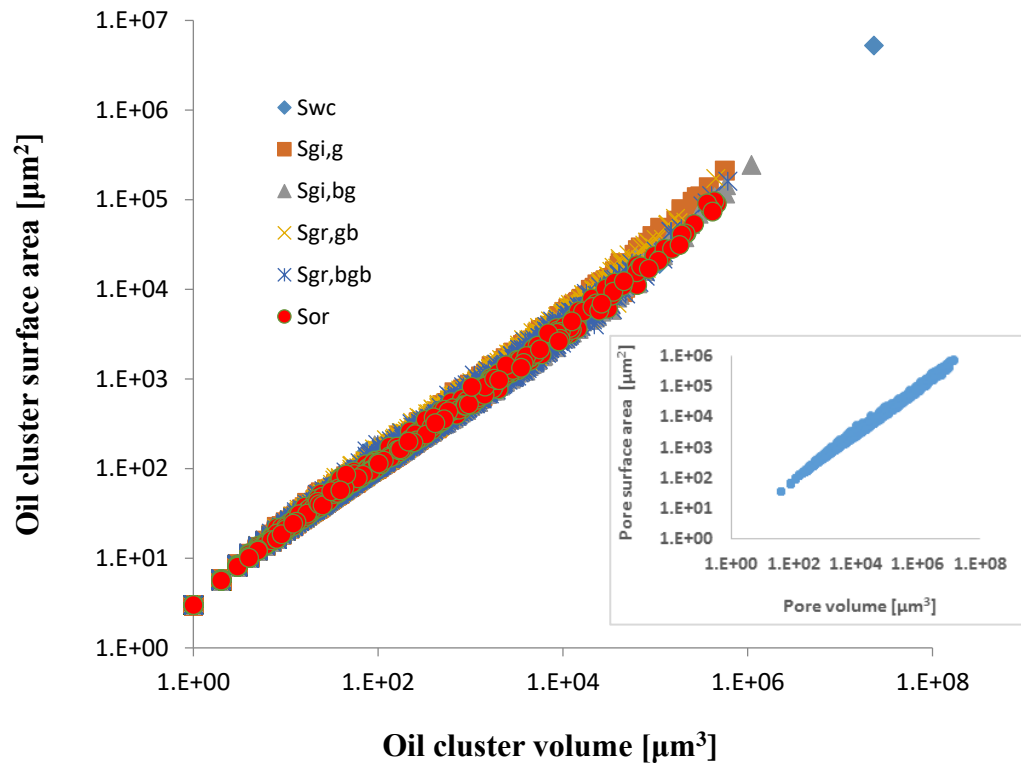


Figure 6-5 Surface area versus cluster volume for each oil cluster and production step.

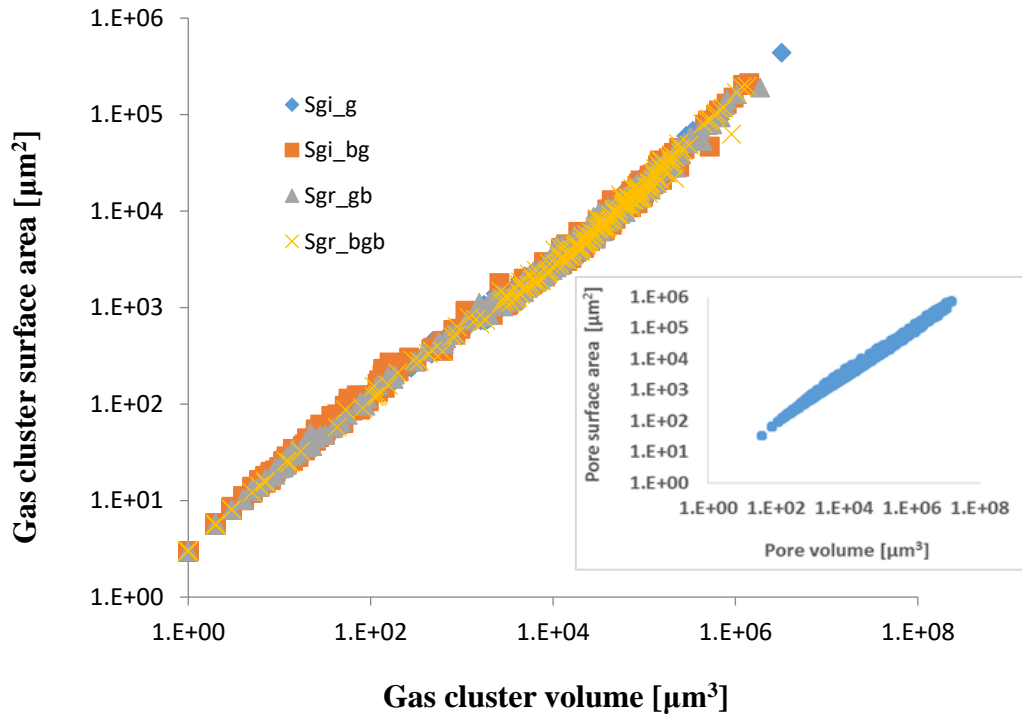


Figure 6-6 Surface area versus cluster volume for each gas cluster and production step.

Table 6-4 Oil and gas cluster statistics: cluster size distribution exponents τ and interfacial area exponents p for intermediate-wet and water-wet sandstone. *Reassessed value: data taken from Iglauer et al. (2014). **Reassessed value: data taken from Iglauer et al. (2013). ***Taken from Iglauer et al. (2013).^{#4}Taken from Pentland et al. (2012).

Saturation state	Intermediate wet		Water-wet		Intermediate wet		Water-wet	
	τ , oil	τ , gas	τ , oil	τ , gas	p , oil	p , gas	p , oil	p , gas
S_{wc}	0.70	---			0.752	---		
S_{or}	0.80	---	0.68*		0.754	---	0.749 ^{#4}	
$S_{gi,bog}$	0.72	0.75			0.765	0.740		
$S_{gi,bobg}$	1.05	0.95			0.754	0.744		
$S_{gr,bogb}$	0.67	1.00	0.9**	0.7**	0.766	0.741	0.729***	0.730***
$S_{gr,bobgb}$	1.00	1.30	1.7**	1.15**	0.756	0.749	0.728***	0.777***

6.3.2.2 Cluster size distributions

Figures 6-7 and 6-8 display the oil and gas cluster size distributions. These distributions can be best fitted with a power-law $N \propto V^{-\tau}$, where N is the frequency with which a cluster of size V is counted in the rock pore space, and τ is a fitting exponent. The best fitting equations are tabulated in Table 6-5. We note that the τ values we obtained (0.7 – 1.7, cp. Table 6-4) are significantly lower than the values estimated previously for two-phase (Iglauer et al. 2010, 2011, 2012, Pentland et al. 2012, Georgiadis et al., 2013) and a three-phase flow scenario (Iglauer et al. 2013). This discrepancy is caused by the statistical methods used previously (Iglauer and Wüling 2015); precisely, previously the methodology suggested by Dias and Wilkinson (1986) was applied, however, this method is biased as it is limited to finding $\tau \geq 2$. We thus used a direct least square fit instead where we qualitatively weighted the points with high counts (i.e. small clusters) more heavily than the points with lower counts, (i.e. larger clusters); this approach is justified as there are only very few large clusters counted, which is caused by the finite size of the observed domain, and the count number for large clusters has thus a lower statistical confidence (cp. Iglauer and Wüling 2015 for a detailed discussion).

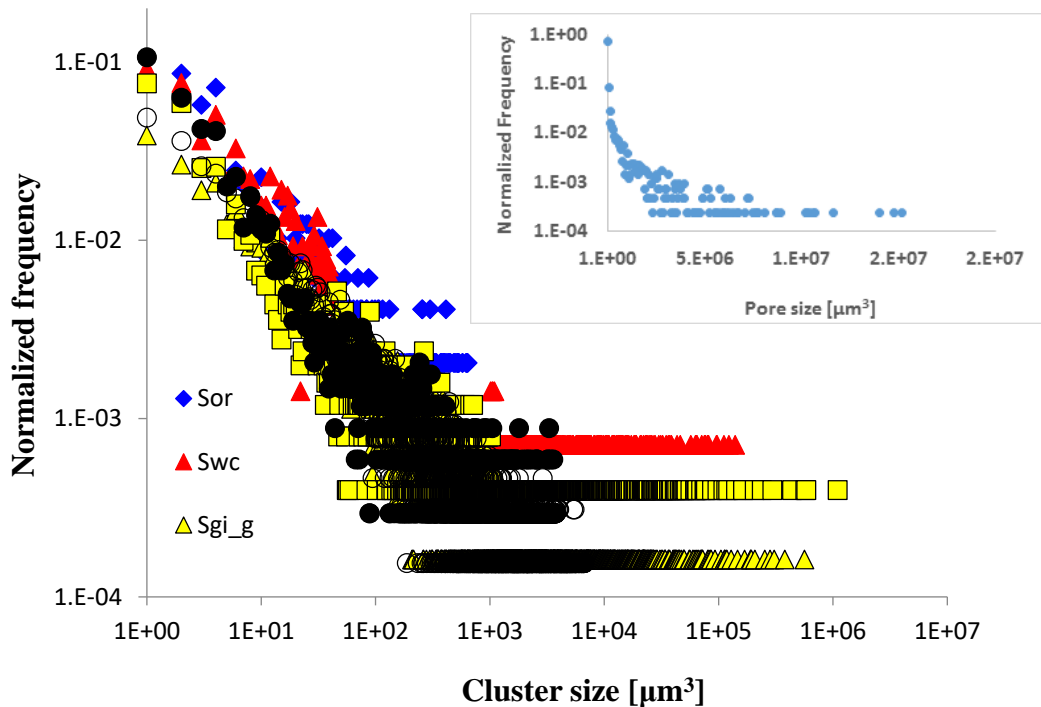


Figure 6-7 Oil cluster size distribution for each saturation state.

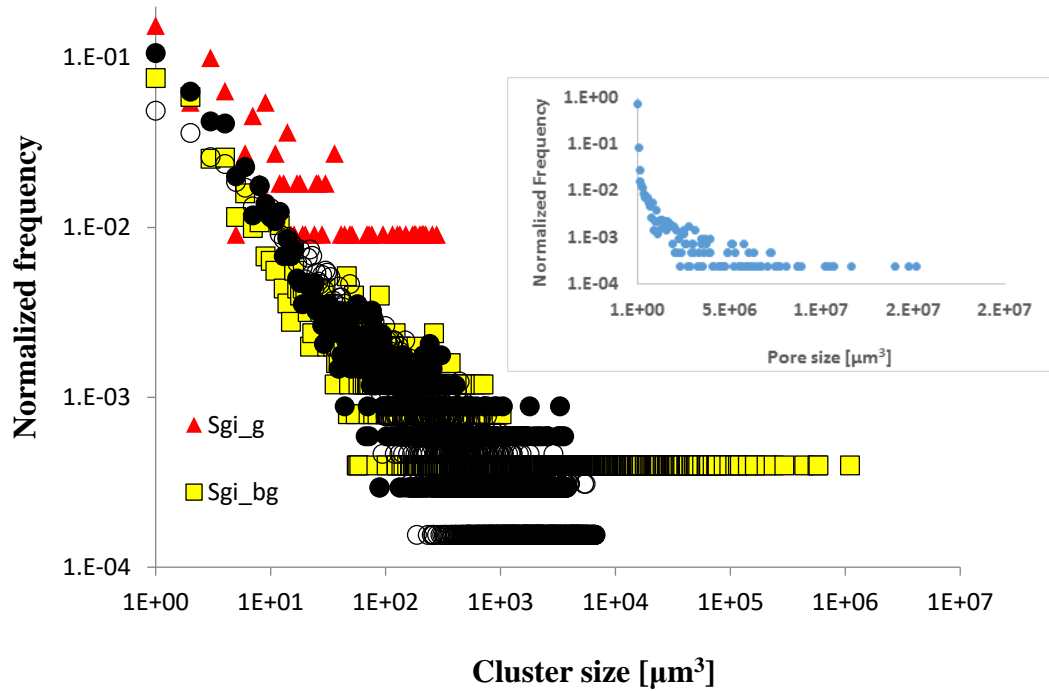


Figure 6-8 Gas cluster size distribution for each saturation state.

Table 6-5 Best least square fitting equations for the measured oil and gas cluster size distributions in water-wet and intermediate-wet sandstone. *Data taken from Iglauer et al. (2014). **Data taken from Iglauer et al. (2013).

Saturation state	Fitting equation			
	Oil	gas	oil	gas
	(intermediate-wet)	(intermediate-wet)	(water-wet)	(water-wet)
S_{wc}	$0.09s^{-0.70}$	---		
S_{or}	$0.11s^{-0.80}$	---	$0.083s^{-0.68*}$	
$S_{gi,bog}$	$0.043s^{-0.72}$	$0.16s^{-0.75}$		
$S_{gi,bogb}$	$0.08s^{-1.05}$	$0.22s^{-0.95}$		
$S_{gr,bogb}$	$0.05s^{-0.67}$	$0.22s^{-1.00}$	$0.16s^{-0.9**}$	$0.136s^{-0.7**}$
$S_{gr,bogb}$	$0.11s^{-1.00}$	$0.38s^{-1.30}$	$0.45s^{-1.7**}$	$0.304s^{-1.15**}$

It is clear that the cluster size distribution curves for the presented cases fall rapidly and there are many more smaller clusters than larger clusters, however, the decay is much smaller than predicted by percolation theory (which predicts $\tau = 2.189$, Lorenz and Ziff 1998). The cluster size distributions again have an important impact on mass transfer processes as they influence interfacial areas, see discussion above and also Figures 6-6 and 6-7, but also re-mobilization of larger drops (which is easier to accomplish, Herring et al. (2013)). The τ values are similar for all saturation states (τ is between ~ 0.7 to 1), except for oil ($\tau = 1.7$) and gas ($\tau = 1.15$) in the water-wet plug and gas ($\tau = 1.3$) in the intermediate-wet plug. The reason for this may be an increased number of snap-off events as gas is always non-wetting and oil is intermediate-wet in a water-wet plug.

6.3.3 Interface curvatures and capillary pressures

6.3.3.1 Interface curvatures

The curvatures (C) of all gas and oil bubbles was measured on smoothed surfaces generated from the segmented images (Armstrong et al. 2012, 2014), Figures 6-9 and 10. For all saturations states the oil curvature distribution was generally quite narrow and always peaked at zero (Figure 6-9) with the exception of S_{wc} , where highly negative curvatures (down to $\sim -40/\mu\text{m}$) were measured, although count numbers were very low. This reflects the intermediate-wet character of the rock (Amott-Harvey index = -0.1, see above). Waterflooding strongly narrowed the distribution and C ranged from ~ -5 to $+5$. Gas and secondary water injection created a tail on the positive side, C increased to up to ~ 20 . These distributions are approximately consistent with those observed for a two-phase oil-water system (Armstrong et al. 2012).

The gas curvature distributions showed a similar behaviour, they also peaked at zero, but a slightly lower symmetry and no tail on the positive curvature side was observed (Figure 6-10).

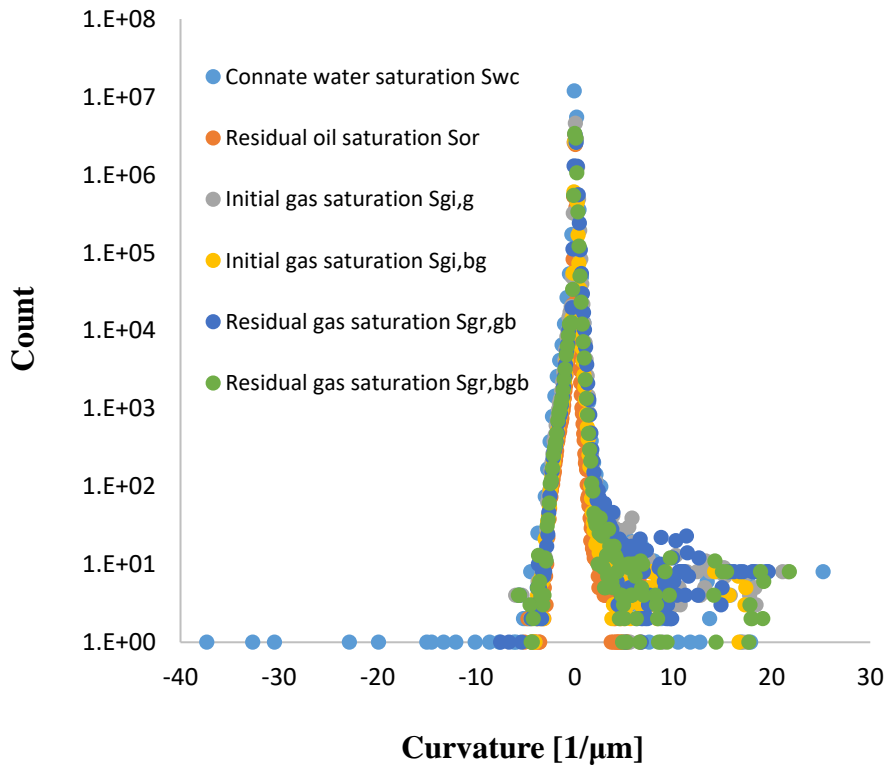


Figure 6-9 Oil curvature distribution at various saturation states.

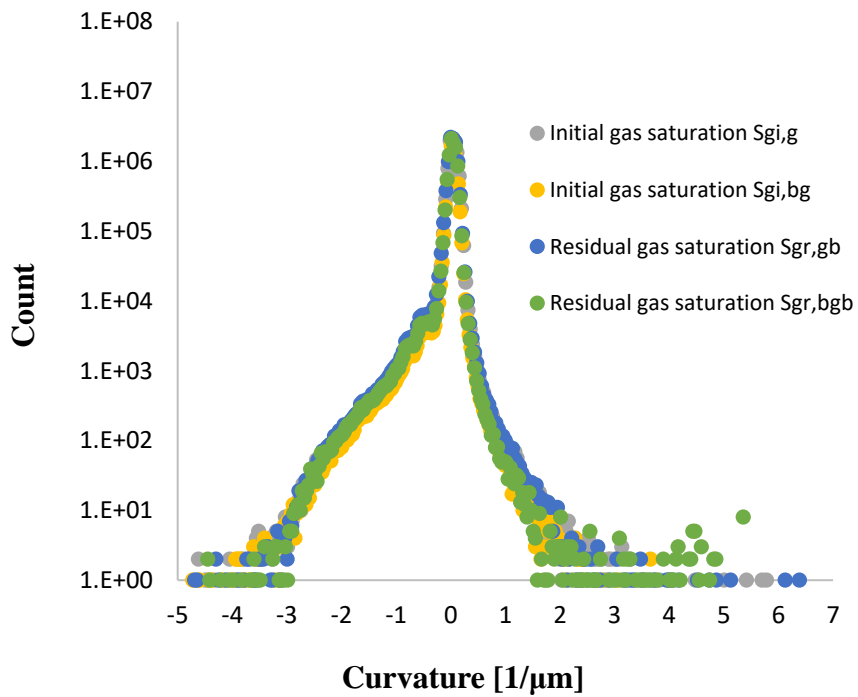


Figure 6-10 Gas curvature distribution at various saturation states.

6.3.3.2 Capillary pressures

The curvatures were then converted into capillary pressures (P_c) using the Young-Laplace equation:

$$P_c = 2\gamma C$$

The resulting capillary pressure distributions of course reflect the curvature distributions, Figures 6-11 and 6-12. Overall remarkably low P_c values were measured at S_{wc} (~ -13 MPa), although count numbers were rare. P_c values for oil were also quite high at the upper end of the tail (after gas and/or secondary waterflood), up to $\sim 7-10$ MPa. This is evidence that intermolecular forces can be very significant at the pore-scale. P_c values were generally lower for gas, although the shape of the curve is similar. The lower numbers were probably due to gas always being the non-wetting phase. We conclude that μ CT is a suitable technique to measure capillary pressures and associated curvatures of individual phase ganglia in-situ.

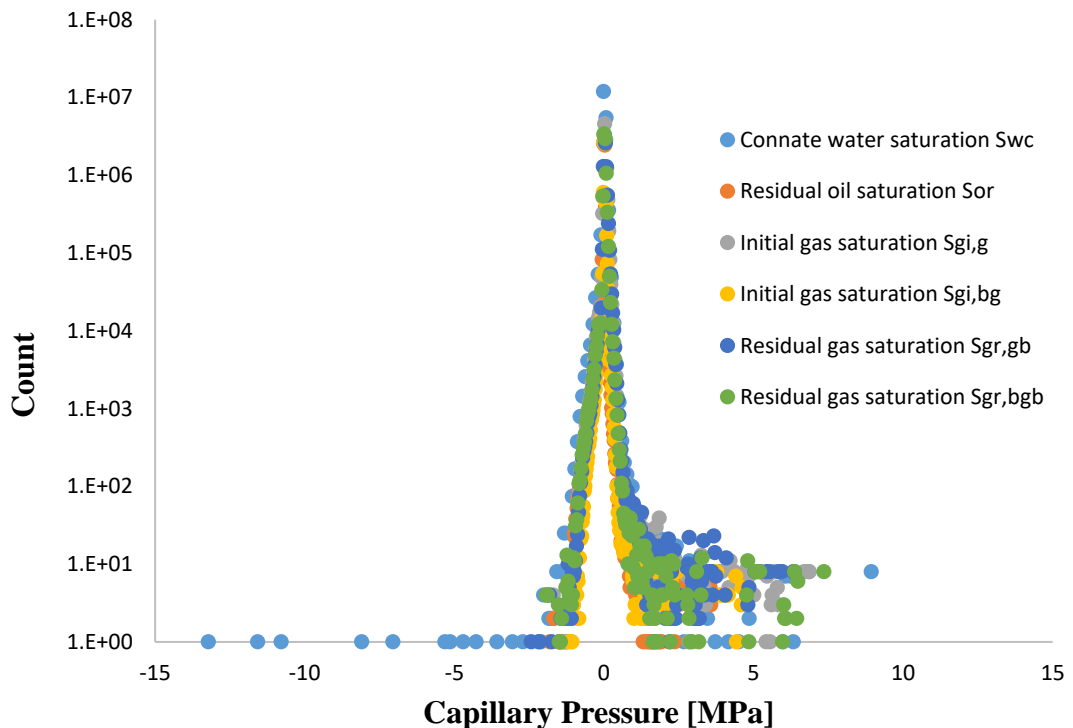


Figure 6-11 Distribution of capillary pressures for oil at various saturation states.

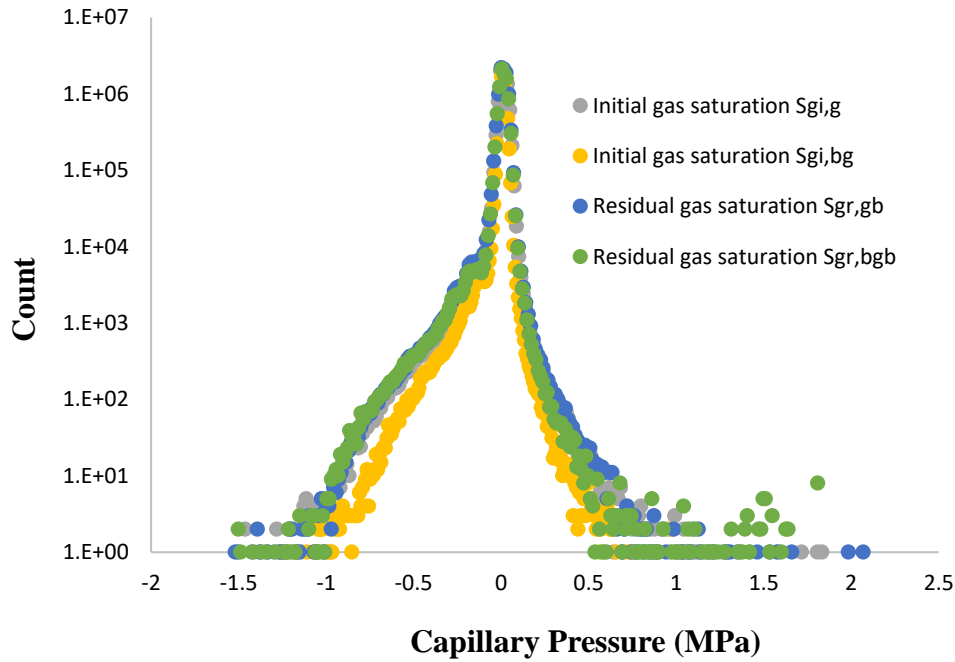


Figure 6-12 Distribution of capillary pressures for gas at various saturation states.

6.4. Conclusions

We imaged two gas-flood assisted enhanced oil recovery processes in an intermediate-wet sandstone at a high resolution of $(3.4\mu\text{m})^3$ in 3D with an x-ray micro-computed tomograph. Waterflooding an intermediate-wet oil reservoir proved to be quite efficient ($R_f \approx 52\%$), although direct gas flooding was even more efficient ($R_f \approx 66\%$). Significant incremental oil could be produced by alternating gas-water flooding ($R_f \approx 71\%$). A large amount of gas ($\sim 50\%$ saturation) could be stored via both processes contrary to what was measured in water-wet sandstone, where $\sim 20\%$ gas could be stored by direct gas flooding and only $\sim 5\%$ was stored when the rock was waterflooded prior to gas injection, Iglauer et al. (2013).

The capillary trapping capacity is thus significantly higher in intermediate-wet rock. The gas and oil cluster size distributions followed power-law relations $N \propto V^{-\tau}$ (where N is the frequency with which clusters of volume V are counted; and the decay exponent τ ranged between 0.7 and 1.7; note that percolation theory predicts a significantly higher τ ($= 2.189$), Lorenz and Ziff 1998). The cluster volume V – cluster surface area A also correlated with a power-law $A \propto V^p$, and p was always \approx

0.75, which is a) significantly smaller than the p predicted by percolation theory (which predicts $p \approx 1$, Stauffer 1979), and b) all values for p measured so far in rocks (Pentland et al. 2012, Iglauer et al. 2013) were ≈ 0.75 for all cases considered.

The fact that significantly lower τ and p values were measured than predicted by percolation theory raises increasing doubt about the applicability of simple percolation models. Furthermore, the curvatures of the oil and gas bubbles were measured and thus the capillary pressure distributions could be determined. The curvatures peaked at zero curvature, consistent with previous measurements in two-phase flow (Armstrong et al. 2012). However, the shapes of the curvature distributions stretched more to positive values in case of oil, and more to negative curvatures in case of gas. These shapes are reflected in the capillary pressure distributions; however, a nominally higher-pressure range was measured for the oil bubbles (from $\sim -13\text{MPa}$ to $+9\text{MPa}$ for oil versus -1.5MPa to $+2.1\text{MPa}$ for gas). Importantly, the detailed pore-scale fluid configurations could be examined at highest scanning resolution $\sim 1 \mu\text{m}^3$; many configurations previously predicted theoretically (Piri and Blunt 2004; Al-Dhahli et al. 2012) or observed in 2D micromodels (Oren and Pincewski 1995) could be verified, and additional, new configurations could be identified; all configurations were categorized in a table.

In conclusion, there is a large variety of possible fluid configurations, fluid curvatures and capillary pressures, and this variety is the origin of the well-known complexity of three-phase flow through intermediate-wet rock.

Chapter 7 Conclusion and Future Work

7.1 Conclusion

Sequestration of carbon dioxide in geological formations is the most direct carbon management strategy for reducing anthropogenic CO₂ in the long term. Further, it has been recognised that residual trapping mechanisms are an effective approach for CO₂ geo-sequestration projects because they immobilise micrometre-sized CO₂ bubbles in the pore network of the rock and increase the safety of the storage sites as they inhibit the upward migration of CO₂. In this study, this mechanism was applied to (oil-wet) oil reservoirs because they are the most important and have been observed for the first time here.

In this thesis, the wettability of CO₂/brine system was measured on hydrophobic and hydrophilic dolomite surfaces at different pressures and temperatures. The experimental observations with the concept of temperature-dependency of van der Waals potential and sharp-kink approximation has been analysed here. It was shown that as pressure increases, the dolomite becomes more CO₂-wet, regardless of the surface being hydrophilic or hydrophobic. The wettability of dolomite, on the other hand, remained unchanged or became slightly more CO₂-wet on hydrophilic surfaces by increasing temperature but was reduced on hydrophobic surfaces. It was illustrated that phase-density difference, fluid interfacial tension, and fluid-solid van der Waals potentials, affect the contact angle (wettability) of no/weakly charged surfaces when state variables, such as pressure and temperature, vary.

Furthermore, intermediate-wet sandstone was imaged in 3D at high resolution (1-3.4 μm)³ with x-ray micro-computed tomography at various saturation states. Initially, the core was at connate water saturation and contained a large amount of oil (94%), which was produced by a waterflood (recovery factor $R_f = 52\%$ OOIP) or a direct gas flood ($R_f = 66\%$ OOIP). Subsequent water and/or gas flooding (WAG process) resulted in significant incremental oil recovery ($R_f = 71\%$ OOIP) while a

substantial amount of gas could be stored (~50%)—significantly more than in an analogue water-wet plug. The oil and gas cluster size distributions were measured and followed power-law correlations $N \propto V^{-\tau}$ (where N is the frequency with which clusters of volume V are counted) with decay exponents τ between 0.7 and 1.7. Further, the cluster volume V plotted against cluster surface area A also correlated with a power-law $A \propto V^p$, and p was always ≈ 0.75 . The measured τ and p values are significantly smaller than predicted by percolation theory (which predicts $p \approx 1$ and $\tau = 2.189$); which raises increasing doubts about the applicability of simple percolation models. In addition, curvatures and capillary pressures of the oil and gas bubbles in-situ were measured and the detailed pore-scale fluid configurations analysed. The complex variations in fluid curvatures, capillary pressures, and the fluid-fluid or fluid-fluid-fluid pore scale configurations (exact spatial locations also in relation to each other and the rock surface) are the origin of the well-known complexity of three-phase flow through rock.

In addition, it was found that the residual CO_2 saturations ($S_{\text{CO}_2,r}$) in oil-wet sandstone are dramatically lower than in equivalent water-wet sandstone ($S_{\text{CO}_2,r} = 8.7\%$ versus 14.9%). This is consistent with data reported for a plastic bead pack (Chaudhary et al. 2013) and indicates that oil-wet rock is intermediate-wet or CO_2 -wet at reservoir conditions—a phenomenon that was also observed in independent contact-angle measurements (Dickson et al. 2006; Yang et al. 2007, 2008; Espinoza et al. 2010; Ameri et al. 2013;). This shift in CO_2 -wettability also significantly influenced CO_2 cluster morphologies and residual cluster size distributions. Flatter and more sheet-like clusters and more small CO_2 bubbles were observed in the oil-wet sample. Therefore, it has been concluded that, importantly, residual trapping is less efficient in oil reservoirs than in deep saline aquifers.

Moreover, the simple experimental cells that allow micro-tomographic visualization of fluids inside rocks under elevated pressure were presented. The modular cell design allows hassle-free maintenance and easy cell modifications. The cell can be used for visualising multiphase fluid distribution inside porous media. Using developed cells, it was investigated how the residual CO_2 phase spreads in the rock as a function of thermo-physical conditions (for ambient condition CO_2 gas, liquid

CO₂, near-critical CO₂ gas, and supercritical CO₂) and how residual CO₂ is dissolved by under-saturated formation brine (dissolution trapping). It was also found that CO₂ spread is relatively homogeneous throughout the core plugs for all conditions. However, less scCO₂ was trapped at higher pressures, which is possible owing to the lower water-wettability of sandstone at high pressures.

7.2 Future Work

While this work has a wide discussion of issues associated with pore-scale investigation in porous media concerning CO₂ geo-sequestration, there are certain other related issues that need to be addressed.

- There is a serious lack of experiments and studies in a three-phase (oil-water-CO₂) system, where oil and water are always present in subsurface formation (meanwhile, the CO₂ is injected for disposal at the pore-scale (micrometre-scale)). Furthermore, determining the whole (hectometre-scale) reservoir behaviour (pressure, temperature, and physical properties) (Bear 1972; Sahimi 2011; Blunt 2013) is even more complicated. This generates a high uncertainty in storage capacity and containment security predictions and, thus, project risk.
- Wettability plays an important role in the underground geological storage of carbon dioxide as the fluid flow and distribution mechanism within the porous medium is controlled by this phenomenon. It is documented that the CO₂ pressure, temperature, brine composition, and minerals type have significant effects on wettability. Despite past research carried out on the subject, the factors controlling the wettability variation for CO₂/water/minerals, particularly the effect of pores of the porous substrate on the contact angle at different pressure, temperature, and salinity and the physical processes, involved are not yet fully understood.
- The contact angle measurement of pure clay minerals (kaolinite, montmorillonite, and illite with CO₂/brine, nitrogen/brine and nitrogen/oil (decane)) at different pressures and temperatures was not systematically studied. Although it was well recognized the temperature, pressure, and salinity have significant effects on contact angle. However, they could considerably differ in subsurface formation. Moreover, it is important to study specifically the wettability in clay minerals, as the materials are also critical to oil recovery in clay-bearing reservoirs and the capacity of capillary seals, and is still not fully understood.

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Zhang, Y., Lebedev, M., Sarmadivaleh, M., Barifcani, A., Rahman, T., Iglauer, S. 2016b. Swelling effect on coal micro structure and associated permeability reduction. Fuel;182:568-576.

Zhang, Y., Xu, X., Lebedev, M., Sarmadivaleh, M., Barifcani, A., Iglauer, S. 2016. Multi-scale x-ray computed tomography analysis of coal microstructure and permeability changes as a function of effective stress. International Journal of Coal Geology;165:149-156

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Appendices

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Al-Yaseri, Roshan, H., A., Zhang, Y., **Rahman, T.**, Barifcani, A., Iglauer, S., 2017. Effect of Temperature on CO₂/Brine/Dolomite Wettability: Hydrophilic vs Hydrophobic Surfaces. Journal of Energy & Fuels.

Statement of Contribution of Others for “Effect of Temperature on CO₂/Brine/Dolomite Wettability: Hydrophilic vs Hydrophobic Surfaces”

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I, Dr. Hamid Roshan, contributed by samples preparation, and measurement to the paper publication entitled

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(Signature of Co-author 1)

Roshan, H



(Signature of First author)

Ahmed Al-Yaseri

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To Whom It May Concern

I, Mr. Yihuai Zhang., contributed by samples preparation to the paper publication entitled

Al-Yaseri, Roshan, H., A., Zhang, Y., **Rahman, T.**, Barifcani, A., Iglauer, S., 2017. Effect of Temperature on CO₂/Brine/Dolomite Wettability: Hydrophilic vs Hydrophobic Surfaces. Journal of Energy & Fuels.

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张艺怀

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Yihuai Zhang

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Ahmed Al-Yaseri

Statement of Contribution of Others for “Effect of Temperature on CO₂/Brine/Dolomite Wettability: Hydrophilic vs Hydrophobic Surfaces”

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Rahman, T



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Ahmed Al-Yaseri

Statement of Contribution of Others for “Effect of Temperature on CO₂/Brine/Dolomite Wettability: Hydrophilic vs Hydrophobic Surfaces”

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Barifcani, A



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Ahmed Al-Yaseri

Statement of Contribution of Others for “Effect of Temperature on CO2/Brine/Dolomite Wettability: Hydrophilic vs Hydrophobic Surfaces”

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I, Prof Stefan Iglauer, contributed by specialist technical advice and manuscript editing to the paper/publication entitled

Al-Yaseri, Roshan, H., A., Zhang, Y., **Rahman, T.**, Barifcani, A., Iglauer, S., 2017. Effect of Temperature on CO2/Brine/Dolomite Wettability: Hydrophilic vs Hydrophobic Surfaces. Journal of Energy & Fuels.

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(Signature of Co-author 5)

Iglauer, S



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Ahmed Al-Yaseri

Appendix 1.2: Statement of Contributions of Others for

M Lebedev, Y Zhang, V Mikhaltsevitch, S Iglauer, **T Rahman**, “Residual trapping of supercritical CO₂: direct pore-scale observations using a low cost pressure cell for micro computer tomography”, Energy Procedia, 114, 4967-4974, 2017

Statement of Contribution of Others for “Residual trapping of supercritical CO₂: direct pore-scale observations using a low cost pressure cell for micro computer tomography”

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张艺怀

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Yihuai Zhang

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Lebedev, M

Statement of Contribution of Others for “Residual trapping of supercritical CO₂: direct pore-scale observations using a low cost pressure cell for micro computer tomography”

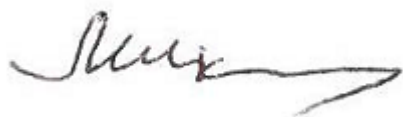
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I, Vassili Mikhaltsevitch, contributed by samples preparation, and measurement to the paper publication entitled

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V, Mikhaltsevitch

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(Signature of First author)

Lebedev, M

Statement of Contribution of Others for “Residual trapping of supercritical CO₂: direct pore-scale observations using a low cost pressure cell for micro computer tomography”

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Iglauer, S

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Lebedev, M

Statement of Contribution of Others for “Residual trapping of supercritical CO₂: direct pore-scale observations using a low cost pressure cell for micro computer tomography”

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I, Taufiqurrahman, contributed by all analysis of residual CO₂ trapping data to the paper publication entitled

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Rahman, T

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Lebedev, M

Appendix 1.3: Statement of Contributions of Others for

Rahman, T., Lebedev, M., Barifcani, A. and Iglauer, S., 2016. Residual trapping of supercritical CO₂ in oil-wet sandstone. *Journal of colloid and interface science*, 469, pp.63-68.

Statement of Contribution of Others for “Residual trapping of supercritical CO₂ in oil-wet sandstone”

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I, Prof Maxim Lebedev, contributed by samples preparation, and measurement to the paper publication entitled

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Lebedev, M



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Rahman, T

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I, Prof Ahmed Barifcani, contributed by specialist technical advice and manuscript editing to the paper/publication entitled

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A handwritten signature in blue ink that reads "A. Barifcani".

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Barifcani, A

A handwritten signature in black ink that reads "Taufiqurrahman".

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Rahman, T

Statement of Contribution of Others for “Residual trapping of supercritical CO₂ in oil-wet sandstone”

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Handwritten signature of Stefan Iglauer, consisting of the letters 'Stefan' followed by a stylized 'Iglauer'.

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Iglauer, S

Handwritten signature of Taufiqurrahman, featuring a stylized 'T' followed by 'aufiqurrahman'.

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Rahman, T

Appendix 1.4: Statement of Contributions of Others for

Iglauer, S., **Rahman, T.**, Sarmadivaleh, M., Al-Hinai, A., Fernø, M.A. and Lebedev, M., 2016. Influence of wettability on residual gas trapping and enhanced oil recovery in three-phase flow: a pore-scale analysis by use of microcomputed tomography. *SPE Journal*.

Statement of Contribution of Others for “Influence of wettability on residual gas trapping and enhanced oil recovery in three-phase flow: a pore-scale analysis by use of microcomputed tomography”

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Sarmadivaleh, M



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Iglauer, S

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Al-Hinai, A

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Iglauer, S

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Fernø, M.A



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Publication: Energy & Fuels

Publisher: American Chemical Society

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