



Tectonic evolution of the southern margin of the Amazonian craton in the late Mesoproterozoic based on field relationships and zircon U-Pb geochronology

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ABSTRACT

New U-Pb zircon geochronological data integrated with field relationships and an airborne geophysical survey suggest that the Nova Brasilândia and Aguapeí belts are part of the same monocyclic, metaigneous and metasedimentary belt formed in the late Mesoproterozoic (1150 Ma-1110 Ma). This geological history is very similar to the within-plate origin of the Sunsás belt, in eastern Bolivia. Thus, we propose that the Nova Brasilândia, Aguapeí and Sunsás belts represent a unique geotectonic unit (here termed the Western Amazon belt) that became amalgamated at the end of the Mesoproterozoic and originated through the reactivation of a paleo-suture (Guaporé suture zone) in an intracontinental rift environment. Therefore, its geological history involves a short, complete Wilson cycle of ca. 40 Ma. Globally, this tectonic evolution may be related with the final breakup of the supercontinent Columbia. Mafic rocks and trondhjemitites in the northernmost portion of the belt yielded U-Pb zircon ages ca. 1110 Ma, which dates the high-grade metamorphism and the closure of the rift. This indicates that the breakup of supercontinent Columbia was followed in short sequence by the assembly of supercontinent Rodinia at ca. 1.1-1.0 Ga and that the Western Amazon belt was formed during the accretion of the Arequipa-Antofalla basement to the Amazonian craton.

Key words: Amazonian craton, intracontinental rift, Mesoproterozoic, Western Amazon belt, U-Pb geochronology.

INTRODUCTION

The Amazonian craton constitutes a central core of the South American continent and it was formed by the amalgamation of Proterozoic accretionary belts that are successively younger toward the southwestern border of the craton (Cordani et al. 2000, Santos et al. 2000, 2008). This continental

landmass has participated in major paleogeographic reconstructions, such as the formation of the Columbia and Rodinia supercontinents (e.g., Rogers and Santosh 2003, Li et al. 2008). The Rodinia supercontinent was formed in the late Mesoproterozoic to early Neoproterozoic by the agglutination of the existing cratonic fragments, such as Amazonia and Laurentia. In Amazonia, the records of this agglutination are present in the

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southwesternmost part of the Amazonian craton. This region is composed of rocks belonging to the Sunsás Orogen, according to Santos et al. (2008), or the Rondonian-San Ignacio Province (Tassinari et al. 1996). The first authors proposed that the Sunsás Orogen is characterized by an autochthonous evolution that lasted for 360 m.y. (from 1465 to 1110 Ma) and was formed by four orogenies: Santa Helena (1460-1420 Ma), Candeias (1370-1320 Ma), San Andrés (1275 Ma) and Nova Brasilândia (1180-1110 Ma). They also postulated the nonexistence of the Paraguá craton, a cratonic fragment that consists of metamorphic basement rocks in the eastern Bolivia Precambrian shield, which was first defined by Litherland et al. (1986), and argued that Paraguá “craton” and Amazonia were already close together, as parts of the same continental masses in the middle to late Mesoproterozoic. In the same way, Sadowski and Bettencourt (1996) argued that there is no clear evidence that the Paraguá craton is a unit that is independent from the Amazonian craton, and they consider it to be part of the Amazonian craton that is defined as the Paraguá Block.

Additionally, Bettencourt et al. (2010), in a recent review of the evolutionary history of the Rondonian-San Ignacio Province, characterized the Province as a composite orogen, consisting of an older, complex accretionary orogen (1556-1430 Ma) that was followed by the collision of the Paraguá Block with the southern part of the cratonized Amazonian craton at 1340-1320 Ma.

The proposal for the tectonic framework of the Amazon made by Tohver et al. (2004, 2005a, b, 2006) draws a comparison between the Grenville history recorded in the SW Amazonian craton with that observed in southernmost Laurentia. They propose that during the late Mesoproterozoic, the SW Amazonian craton was marked by two separate tectonic events, the first related to collision with southern Laurentia (ca. 1.2-1.15 Ga) and the second caused by suturing of the Paraguá craton

(ca. 1.09 Ga). These researchers also suggest the existence of a suture belt (Nova Brasilândia belt) indicating that the initial docking of the Amazonian craton with southern Laurentia (Llano uplift) at 1.2 Ga was followed by strike-slip motion, allowing for later suturing of the Paraguá craton in the final stages of the amalgamation of Rodinia. They also suggest that the Nova Brasilândia sedimentation occurred during a transtensional phase at < 1.2 Ga.

In contrast, alternative positions of the Amazonian craton in Mesoproterozoic are envisaged to accommodate smaller crustal fragments, such as the Arequipa-Antofalla (i.e., Keppie and Ortega-Gutierrez 1999, Loewy et al. 2003, 2004, Casquet et al. 2006). In this way, Boger et al. (2005) present three possible scenarios for the accretion of the Paraguá Block and the Arequipa-Antofalla basement to the western margin of the Amazonian craton. Nevertheless, latest authors favor the tectonic model by which the Paraguá Block was accreted to the margin of a coherent Amazonian craton during the Sunsás Orogeny.

The tectonic framework of the Amazon basement (Rondônia-Juruena Province) is of a multi-orogenic nature, whereas the adjacent Nova Brasilândia Group in the Nova Brasilândia belt most likely originated in an intracontinental rift setting. Rifting was followed by the opening of a proto-ocean at <1215 Ma as part of a single Wilson cycle (Rizzotto 1999).

With a similar geochronological and tectonic evolution, the Sunsás and Aguapeí belts occur along the southeastern rim of the Paraguá Block. These belts are most likely related to the timing of an intraplate extension episode, a stage equivalent to the Nova Brasilândia belt passive margin. The deposition age of the Aguapeí Group occurred between 1167 ± 27 Ma and 1149 ± 7 Ma (Santos et al. 2005). However, Tohver et al. (2004, 2005a) and Teixeira et al. (2010) suggest the absence of physical and structural continuity between the northernmost extent of the Aguapeí belt and the

E–W trending Nova Brasilândia belt. Teixeira et al. (2010) suggest that the Aguapeí belt is a late tectonic component of the Sunsás province as a result of the Sunsás collision.

The stratigraphic correlation between the Aguapeí and Sunsás Group is in agreement with the idea that late Mesoproterozoic extensional dynamics produced a system of continental, aborted rifts over the cratonized SW margin of the Amazonian craton, evolving to a wide intracratonic basin (e.g., Leite and Saes 2003).

Therefore, the inconsistencies regarding the collided Paraguá Block–Amazonian craton and its geodynamic behavior during the Mesoproterozoic are under evaluation. However, Rizzotto and Hartmann (2012) suggested a new scenario for the operation of Mesoproterozoic plate tectonics, with the definition of the Guaporé suture zone, a mega-structure that marks the suturing of the proto–Amazonian craton and the Paraguá Block in the late Calymmian to middle Ectasian, ca. 1470–1320 Ma. The Guaporé suture is a zone of crustal weakness reactivated during sedimentation and the proto-oceanic opening of the Nova Brasilândia basin.

In this paper, we present new zircon U–Pb geochronological results from the two Mesoproterozoic Nova Brasilândia and Aguapeí orogenies of the southwestern Amazonian craton, coupled with field relationships, in a contribution to resolved some of these issues. These results place new delimitations on the geologic evolution of the southwestern margin of the Amazonian craton, and this sheds light on paleogeographic correlations between Amazonia and Laurentia.

GEOLOGICAL OVERVIEW

The studied area is located between the Sunsás and Rondônia–Juruena Provinces (Santos et al. 2008), southwestern Amazonian craton (Fig. 1). The basement rocks of the Rondônia–Juruena Province are included in six major Paleo–Mesoproterozoic terranes and events: the Juruena magmatic arc,

the Jamari magmatic arc, the Alto Juru Terrane, the Quatro Cachoeiras orogeny, the Cachoeirinha orogen and the Anorogenic magmatism (Fig. 1).

The Paleoproterozoic basement rocks of the Sunsás Province (Paraguá Block) (Fig. 1) comprise three main units: the Lomas Manechi Complex display granulitic metamorphism generated during the Mesoproterozoic, at ca. 1330–1350 Ma (Santos et al. 2008). The Chiquitania Complex and San Ignacio Group subsequently deformed and metamorphosed during the San Ignacio Orogeny by means of three successive events of deformation accompanied by voluminous syn-tectonic granite magmatism (Litherland et al. 1986).

This Paleoproterozoic basement was amalgamated along WNW-trending Mesoproterozoic belts, described here from oldest to youngest. The Alto Guaporé belt (Rizzotto and Dehler 2007) is a accretionary orogen formed by oceanic plate subduction between 1.47–1.43 Ga (Rizzotto and Hartmann 2012). This constitutes the Guaporé suture zone between the proto–Amazonian craton and the Paraguá Block. The rocks of the Alto Guaporé belt are exposed from the northwestern sector of Rondônia state to the southeastern and to the southwestern portion of Mato Grosso state. They are largely covered by the Cenozoic sedimentary rocks of the Parecis and Guaporé Basins, in an area of 200 x 60 km, with WNW–ESE trend marked by strong magnetic anomalies. This belt consists predominantly of, plutonic, sedimentary and mafic-ultramafic volcanic rocks of the Trincheira Complex and calc-alkaline granitic rocks. The Colorado Complex is a meta-sedimentary sequence occurring in the Guaporé suture and is composed of paragneiss, pelitic schist, calc-silicate gneiss, para-amphibolite and banded iron formation. The protolith of the paragneiss and pelitic schist is interpreted as a turbiditic sequence deposited in a passive margin (Rizzotto and Quadros 2007).

The Nova Brasilândia and Aguapeí belts (ca. 1150–950 Ma) are exposed along the Bolivia–Brazil border (Fig. 2). The Nova Brasilândia belt

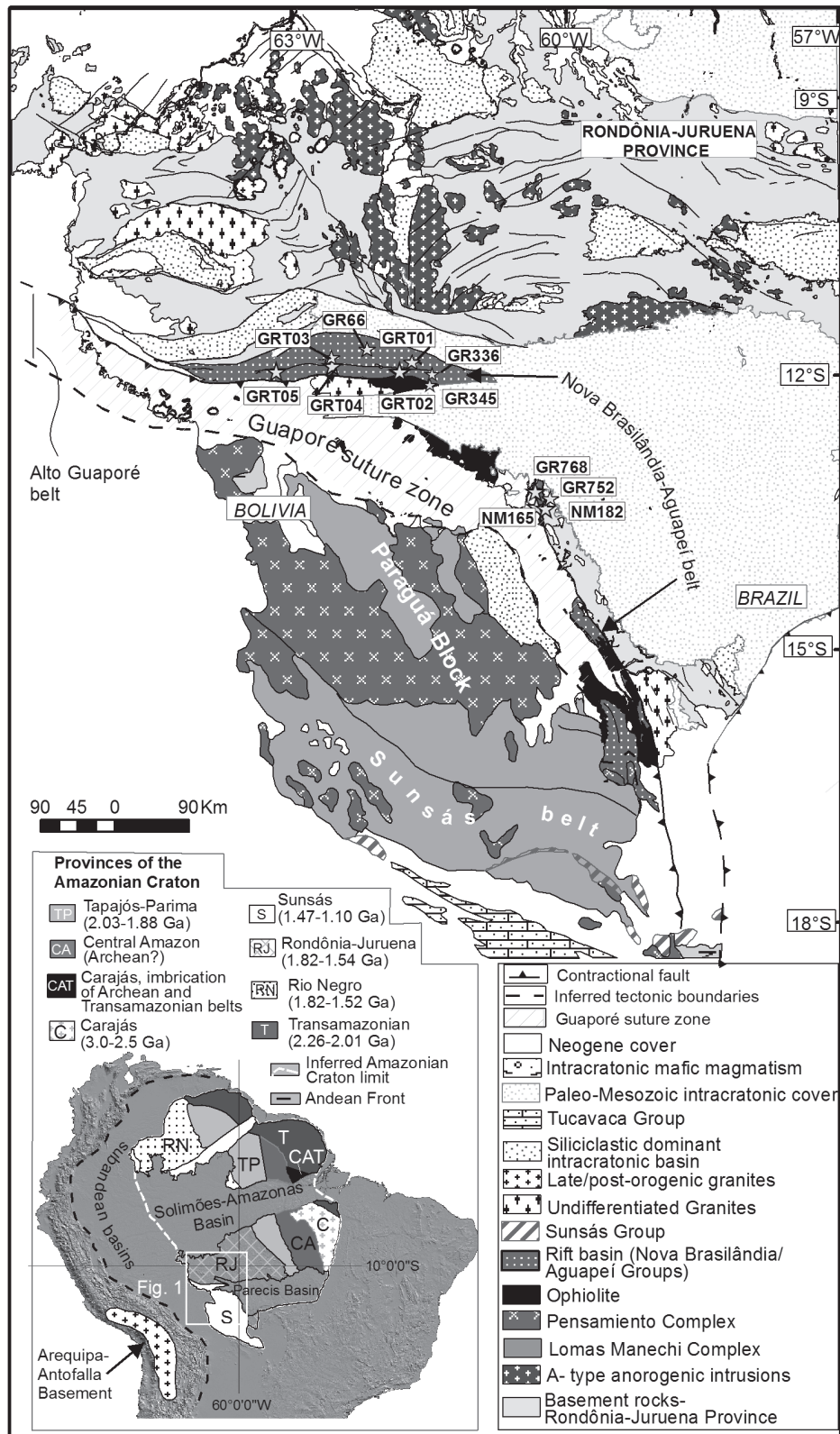


Figure 1 - Simplified map of the southwestern Amazonian craton showing the approximate boundaries of the main Provinces, belts, tectonic elements, lithologic units, and location of dated samples.

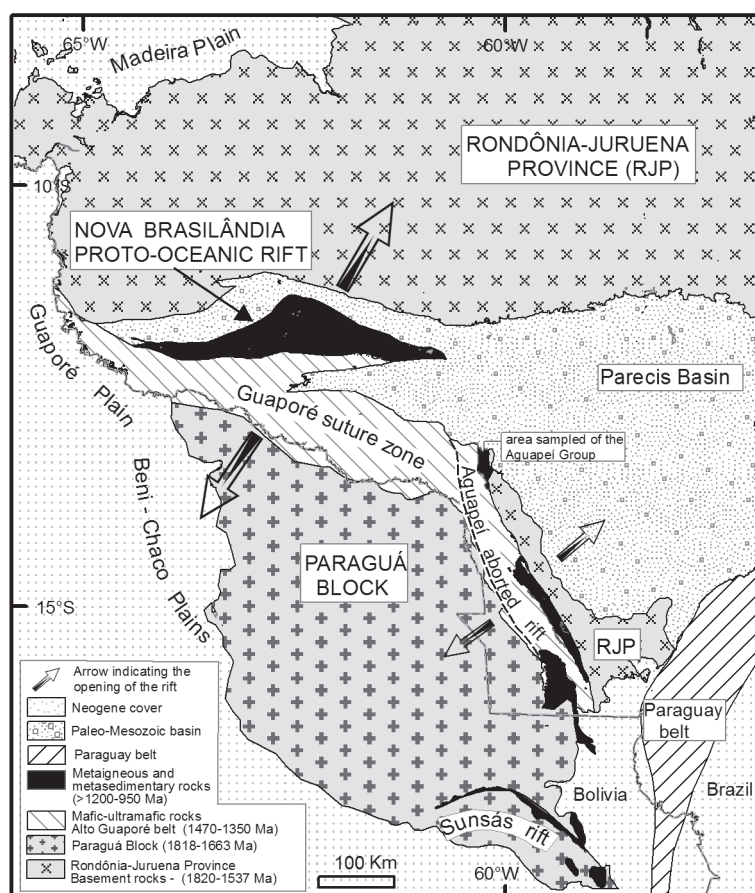


Figure 2 - Geotectonic map of the southwestern Amazonian craton highlighting the Nova Brasilândia, Aguapeí and Sunsás rifts.

originated from the opening and closing of an ocean, whereas the Aguapeí belt is an aborted rift, both generated by reactivation along the Guaporé suture zone at ca. 1110-950 Ma.

The region being examined comprises the Nova Brasilândia and Aguapeí belts, with the integration of new field and geochronological data. The Sunsás belt is also described for comparison. The importance of the theme is centered on the register of the last orogenic phase in the southwestern border of the Amazonian craton and its correlation with the Grenville province in southern Laurentia.

Grenvillian-age deformation in the basement of the Rondônia-Juruena Province

The timing of the deformation and generation of orogenic rocks of Stenian age in the southwestern

Amazonian craton was first estimated by Rizzotto (1999) using U-Pb geochronology in the Nova Brasilândia belt. To the north of this belt, the granitoid basement rocks of the Rondônia-Juruena Province record a polycyclic evolution characterized by the complex juxtaposition of distinct units and widespread shear zone structures.

The granitoid basement rocks were sheared by a wide network of sinistral, amphibolite-grade, strike-slip shear zones defined as the Ji-Paraná shear zone (Scandolaro et al. 1999) which extends over hundreds of kilometers and cross-cut all preexisting rock fabrics. According to Tohver et al. (2005a), metamorphism along the shear zone occurred at temperatures of 450-550°C, and the timing of this deformation varies from ca. 1.18 to 1.15 Ga. Consistent with these findings, recent data

from field work in central Rondônia by Quadros et al. (2012) defined the Ji-Paraná-Cujubim transpressional sinistral system as extending to the northwest of the Ji-Paraná shear zone. Within this zone, the Igarapé Quinze Formation consists of paragneisses and schists that trend N40W. The dating of detrital zircon from the schist yielded U-Pb ages in the range of 1940 Ma and 1367 Ma. The younger age is interpreted as the maximum age of sedimentation of the sedimentary protoliths of the Igarapé Quinze Formation (Quadros et al. 2012). This formation is interpreted as turbiditic deposits and the tectonic structures allow its correlation with the Nova Brasilândia Group (Quadros et al. 2012).

According to these authors, lenses and pods of anatectic granite were generated by partial melting of metapelites. Zircon grains from a locally-derived anatectic melt yield a U-Pb age of 1138 ± 8 Ma. This age represents the peak metamorphic conditions experienced by the sedimentary protolith along the Ji-Paraná-Cujubim belt.

NOVA BRASILÂNDIA BELT

The Nova Brasilândia Group in the Nova Brasilândia belt consisted originally of a supracrustal unit of deep-sea turbidites with a siliceous-clastic-to-carbonatic composition. The unit was metamorphosed in upper-amphibolite to granulite facies to produce quartz-feldspar gneisses, mica schists and minor calc-silicate gneisses. The metasedimentary package also includes metagabbro sills, amphibolites and scarce metabasalts. SHRIMP U-Pb analyses of detrital zircon grains from the Nova Brasilândia pelitic rocks yield ages that range from ca. 2090 ± 17 Ma to 1122 ± 12 Ma. The main population of zircon has an age of 1215 Ma, interpreted by Rizzotto (1999) as the maximum deposition age for the sequence. Sample reprocessing shows a younger group of zircon grains with an age of 1122 ± 12 Ma, interpreted as the age of the basin opening (see section 4).

The other unit consists of mafic sills that outcrop in a narrow and elongated WNW-ESSE belt

of voluminous plutonic mafic rocks with a strong, positive magnetic signature (Fig. 3). This magmatism is related to the successful initial rifting that formed a narrow proto-ocean, currently with approximately 12 x 110 km, during the evolution of continental to oceanic rift phase. These mafic rocks are here considered to record an extensional anorogenic magmatism associated with fragmentation of a pre-Rodinia supercontinent.

This mafic unit is bordered in the north by a transcurrent sinistral shear zone, which consists of mafic granulites (norite and gabbro as protoliths), rare acid granulites, banded amphibolites, porphyroblastic amphibolites, and interbedded trondhjemites, diabase dikes and lenses of calc-silicate gneisses. Rare and scattered basalts, metasedimentary rocks and banded iron formation indicate a high degree of exhumation of the terrain with the accompanying erosion of the top portion of the volcano-sedimentary rocks. The Rio Branco monzogranitic orthogneiss was emplaced under a transpressive regime in the banded amphibolite at ca. 1113 Ma, with $T_{DM}=1630$ Ma (Rizzotto 1999).

The transtensional stage was succeeded by a transpressional regime accompanied by crustal shortening, as evidenced by regional EW- and WNW-ESE-trending structures, dominated by oblique thrust faults (D1). The northeast vergence and medium-to high-grade metamorphism correspond to a progressive thrust to transcurrent shear zones with tectonic transport from SSW to NNE. The main extensional structures are oriented in the N45°E direction, while the structures with a direction of N45°W correspond to a frontal thrust. The synkinematic emplacement of the Rio Branco granite and anatectic leucogranite (1110 ± 8 Ma) was contemporaneous with this compressional stage. These are the first Grenvillian-age orogenic rocks within the Nova Brasilândia belt described by Rizzotto (1999) within the southwestern margin of the Amazonian craton.

The geological history of the Nova Brasilândia belt is supplemented with the emplacement of the

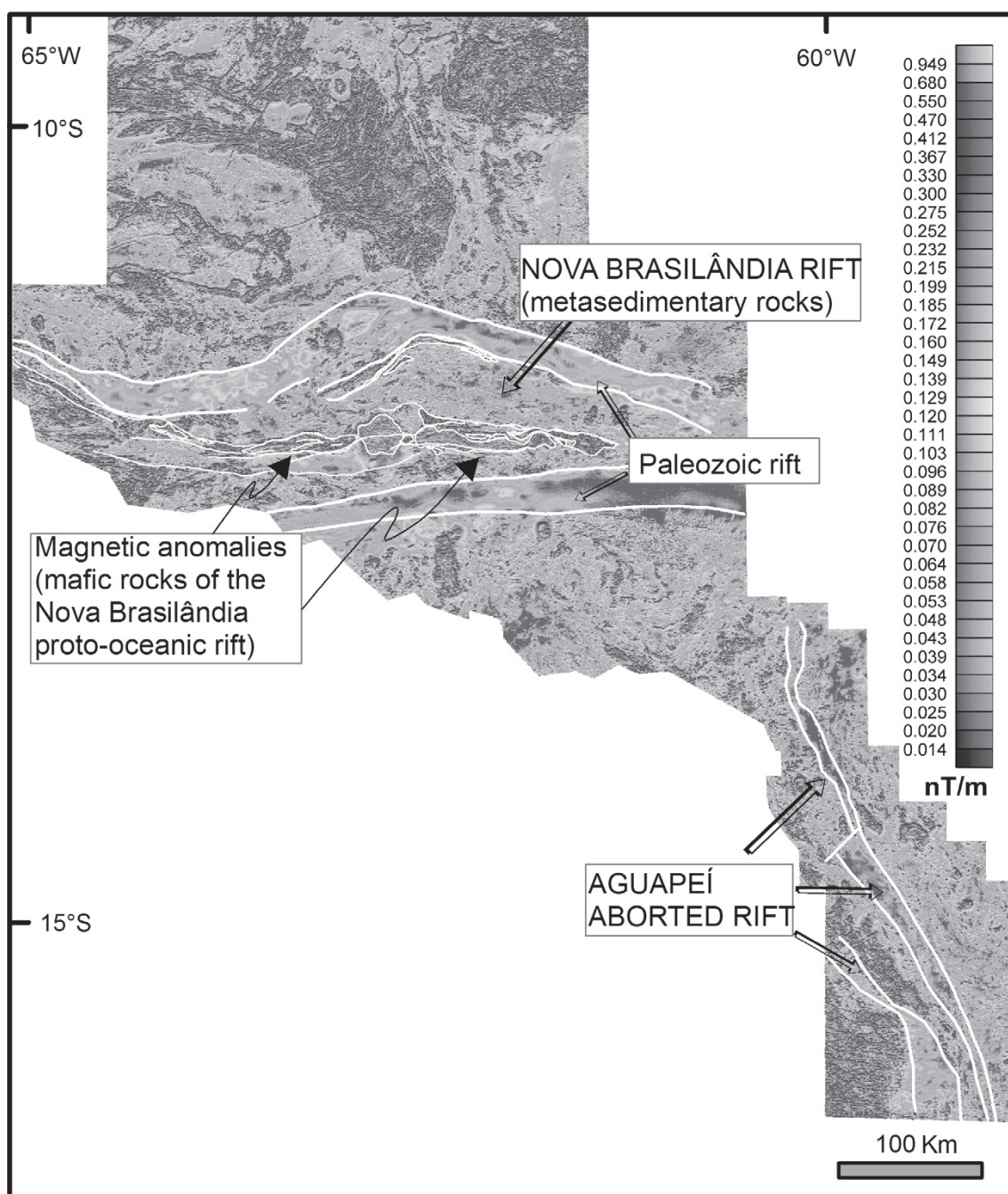


Figure 3 - Magnetometric data (analytical signal amplitude map) of southwestern margin of the Amazonian craton. The highlighted (white line) magnetic lineaments indicate the rift zones and areas with mafic rocks outcrops of the Nova Brasilândia proto-oceanic rift phase.

Rio Pardo granite, which comprises a kilometer-sized syn- to late transcurrent body with an age of the 1005 ± 41 Ma, injected during the development of transcurrent shear zones.

Thus, the Nova Brasilândia Group originated in an intracontinental rift setting followed by the opening of a proto-ocean and was further subjected to crustal shortening, deformation and associated

magmatism. This is suggested by detailed structural studies and metamorphic petrology coupled with chemical and petrogenetic constraints determined from the sedimentary and igneous rocks (Rizzotto 1999). This scenario is supported by recent paleomagnetic results reported by D'Agrella-Filho et al. (2012) where the 1420 Ma Indiavaí and Nova Guarita poles suggest that the Nova Brasilândia belt does not represent a suture zone, but rather an intracratonic structure.

In contrast, Tohver et al. (2004, 2005a) interpreted the Nova Brasilândia belt as the result of the crustal thickening through imbrication caused by the transpressive suturing of the Amazonian and Paraguá cratons at ca. 1.09 Ga. They used detailed structural data, supported by thermobarometric calculations and geochronological information. Metamorphic ages (U-Pb in titanite and monazite; $^{40}\text{Ar}/^{39}\text{Ar}$ in amphibole and biotite) range from 1.09-1.06 Ga to 0.97- 0.91 Ga throughout the belt (e.g., Tohver et al. 2004, 2006). This age range reveals a long history of metamorphism, tectonic exhumation and cooling.

According to Rizzotto (2001), the distensional phase and final orogenic collapse (1005-970 Ma) of the Nova Brasilândia belt was accompanied by late to post-tectonic granites dated at 1.05 Ga ($\epsilon\text{Nd}(T) = +0.5$), with the generation of the foreland basins (Palmeiral Formation) in a cratonic stabilization phase accompanied by bimodal intraplate magmatism. This stage was accompanied by lateral movements of crustal blocks that generated large strike-slip zones (Rio Branco shear zone). This is in agreement with the $^{40}\text{Ar}/^{39}\text{Ar}$ and K-Ar analyses on amphibole from amphibolite (Tohver et al. 2004) and muscovite from granitic mylonite (Tassinari 1993), respectively, along this strike-slip zone. These analyses yielded a minimum age of metamorphism and deformation (D_2) at 970-966 Ma and 965 ± 23 Ma. Similar ages were obtained by Fernandes et al. (2006) in sericite of the metarenites and metaconglomerates of the Aguapeí Group.

AGUAPEÍ BELT

The Aguapeí belt consists of a narrow metasedimentary belt that borders the eastern margin of the Paraguá Block in the southwestern Mato Grosso state. It comprises a folded metasedimentary unit (Aguapeí Group) deposited in an NNW-SSE trending aulacogen (Saes 1999) or in a rift environment (Saes et al. 1992). Souza and Hildred (1980) characterized the Aguapeí Group as a thick sequence of siliciclastic rocks composed, from bottom to top, of the Fortuna, Vale da Promissão and Morro Cristalina formations. The Fortuna Formation shows thick packages of oligomictic conglomerates and quartz sandstones with interbedded metasiltite/metapelite deposited in a shallow sea environment. The Vale da Promissão Formation displays a transitional interfingering contact with the unit underlying and comprises a sequence dominated by metapsamites, metapelites and rare metarenites deposited in a deep marine environment. The Morro Cristalina Formation is composed of fluvial sandstones and siltstones. The depositional age of the Aguapeí Group is estimated between 1167 ± 27 Ma and 1149 ± 7 Ma (Santos et al. 2005).

The Aguapeí Group overlies the Rio Alegre plutono-volcanic rocks, occurring discontinuously along more than 500 km between the northeastern/eastern margins of the Paraguá Block. The depocenter of the Aguapeí Group coincides with the paleo-suture previously established in the Rio Alegre terrane (Guaporé suture zone-Rizzotto and Hartmann 2012). The pronounced magnetic low that characterizes the Aguapeí Group may be partially related to a failed rift that is filled with Mesoproterozoic sediments (Fig. 3). We suggest that the tectonic evolution of the Aguapeí Group was initiated by rifting that subsequently failed, allowing the development of a long-lived intracratonic sedimentary basin.

The Aguapeí Group exhibits pervasive deformation and low-grade metamorphism (e.g., Geraldés et al. 2001, Ruiz 2005) within a regional synform structure indicated by gentle folds

(upright to northeast-dipping) and NW-trending shears. Sinistral shearing and thrusts deformed this sequence. Recrystallized mylonites yielded K/Ar and $^{40}\text{Ar}/^{39}\text{Ar}$ mica ages between 960 and 910 Ma (e.g., Geraldès et al. 1997, Fernandes et al. 2006, Ruiz et al. 2007). Because of the occurrence side by side of folded and sheared rocks (Sunsás and Aguapeí Groups) and flat-lying and unmetamorphosed rocks (Huanchaca and Ricardo Franco ridges), Santos et al. (2002) prefer the usage of Aguapeí I for the folded and metamorphosed unit and Aguapeí II for the flat-lying cover.

SUNSÁS BELT

The sedimentary units of the Sunsás Group were deposited between 1300 and 1000 Ma and unconformably overlie the metamorphic rocks of the basement. They were deformed together with the basement during the subsequent Sunsás Orogeny at ca. 1000 Ma (Litherland et al. 1986).

The Sunsás Group comprises a sedimentary sequence up to 6 km thick, found mainly in the Serranía Sunsás, Bolivia (Litherland and Bloomfield 1981). In general, this group comprises a basal conglomerate overlain by psammitic units (arkoses, sandstones, quartzites), pelitic units (mudstones and siltstones), and a combination of feldspathic sandstones, pelites and quartzites (Berrangé and Litherland 1982). Sedimentary structures, such as cross-beds and ripple marks, are common in the psammitic units, mainly in quartzites of the Serranía Los Tajibos. Berrangé and Litherland (1982) suggested that the sedimentary rocks of the Sunsás Group were deposited in an alluvial to deltaic environment. Later, Saes et al. (1992) referred to the sediments as deposited in an intra-continental rift.

The upper psammitic units (Vibosi Group) of the Sunsás Group were intruded by the layered Rincón del Tigre igneous complex (Prendergast 2000), which yielded a Rb–Sr age of ca. 990 Ma. This age is similar to that obtained for the late syn- to post-tectonic

Sunsás granites (Litherland et al. 1986).

The Sunsás belt is structurally marked by upright NW-trending folds formed contemporaneously with extensive mylonitic shear zones that developed during low-to-medium grade metamorphism, active at ca. 1080–1050 Ma (Boger et al. 2005).

MATERIALS AND METHODS

The samples were crushed, milled and split into fractions for whole rock isotope geochemistry, zircon, rutile and titanite dating. Zircon, titanite and rutile were separated from twelve samples using heavy liquid and magnetic separation techniques. All rocks were investigated in thin section, and their chemical compositions were determined. Nine rock samples were analyzed by SHRIMP at the Curtin University of Technology-Australia, and three samples were analyzed by LA-MC-ICP-MS. All grains used for zircon, titanite and rutile dating were imaged with backscattered electrons to determine their internal structure and the nature (igneous or metamorphic) of their growth zones. The sampling sites are shown in Figure 1, and general data for each sample (coordinates, rock name, and stratigraphic unit) are described in the next item.

The analytical procedures are described in the online edition (Appendix SA) (Supplementary material), and the isotopic data are presented in Table I.

RESULTS

NOVA BRASILÂNDIA BELT

The samples were collected along an EW belt that contains primarily mafic rocks belonging to proto-ocean, the Nova Brasilândia rift. These lithotectonic units were deformed by NW-SE (D_1) to E-W (D_2) transpressional phases, corresponding to a progressive thrust to transcurrent shear zones. Generally, all localities preserve evidence of granulite facies metamorphism associated with ductile deformation; orthopyroxene often

TABLE I
SHRIMP and LA-MC-ICP-MS zircon and titanite U-Pb analytical results for the samples studied.

spot	U ppm	Th ppm	Th U	$^{206}\text{Pb}/^{206}\text{Pb}$	$4f^{206}$ %	Ratios				Ages		Disc. %	
						$^{207}\text{Pb}/^{206}\text{Pb}$	$^{207}\text{Pb}/^{235}\text{U}$	$^{206}\text{Pb}/^{238}\text{U}$	err. corr.	$^{208}\text{Pb}/^{232}\text{Th}$	$^{206}\text{Pb}/^{238}\text{U}$		$^{206}\text{Pb}/^{207}\text{Pb}$
GRT1, zircon													
h.1-1	706	254	0.37	105.7	0.10	0.07548 ± 0.88	1.8118 ± 1.68	0.1741 ± 1.43	0.851	0.0523 ± 1.76	1035 ± 14	1081 ± 18	4
h.1-2	1056	821	0.80	159.1	-0.01	0.07696 ± 0.50	1.8613 ± 1.48	0.1754 ± 1.40	0.942	0.0512 ± 1.48	1042 ± 13	1120 ± 10	7
h.1-3	2254	501	0.23	319.2	0.00	0.07301 ± 0.32	1.6593 ± 1.50	0.1648 ± 1.47	0.978	0.0504 ± 2.81	984 ± 13	1014 ± 6	3
h.2-1	2021	2266	1.16	322.8	0.00	0.07667 ± 0.63	1.9654 ± 1.59	0.1859 ± 1.47	0.920	0.0551 ± 1.60	1099 ± 15	1113 ± 12	1
h.3-1	779	426	0.57	125.8	0.00	0.07669 ± 0.47	1.9886 ± 1.56	0.1881 ± 1.49	0.954	0.0551 ± 1.58	1111 ± 15	1113 ± 9	0
GRT2, titanite													
a.1-1	254	244	0.99	35.7	0.02	0.07229 ± 0.65	1.6302 ± 1.28	0.1636 ± 1.10	0.860	0.2735 ± 1.23	976 ± 10	994 ± 13	2
a.1-2	268	260	1.01	37.9	0.08	0.07152 ± 0.72	1.6231 ± 1.30	0.1646 ± 1.09	0.834	0.1574 ± 2.42	982 ± 10	972 ± 15	-1
a.2-1	234	225	0.99	33.1	-0.05	0.07218 ± 0.74	1.6358 ± 1.32	0.1644 ± 1.09	0.830	0.1723 ± 1.21	981 ± 10	991 ± 15	1
a.3-1	186	180	1.00	26.8	0.13	0.07191 ± 0.68	1.6553 ± 1.30	0.1670 ± 1.11	0.852	0.1672 ± 1.25	995 ± 10	983 ± 14	-1
a.3-2	215	211	1.02	30.8	0.00	0.07231 ± 0.80	1.6619 ± 1.41	0.1667 ± 1.16	0.823	0.1663 ± 1.29	994 ± 11	995 ± 16	0
GRT2, zircon													
a.5-1	509	66	0.13	80.2	0.01	0.07363 ± 0.86	1.8597 ± 1.70	0.1832 ± 1.47	0.864	0.0537 ± 2.55	1084 ± 15	1031 ± 17	-5
a.5-2	112	40	0.37	18.8	0.07	0.07571 ± 1.57	2.0354 ± 2.34	0.1950 ± 1.73	0.740	0.0590 ± 2.79	1148 ± 18	1087 ± 31	-6
a.9-1	1201	40	0.03	191.5	-0.01	0.07537 ± 0.52	1.9295 ± 1.54	0.1857 ± 1.44	0.940	0.0581 ± 2.42	1098 ± 15	1078 ± 11	-2
a.10-1	141	50	0.37	23.3	0.00	0.07575 ± 1.72	2.0036 ± 2.40	0.1918 ± 1.68	0.699	0.0575 ± 2.51	1131 ± 17	1089 ± 34	-4
a.10-2	398	202	0.53	65.1	0.06	0.07606 ± 0.89	1.9973 ± 1.74	0.1905 ± 1.50	0.859	0.0590 ± 1.82	1124 ± 15	1097 ± 18	-2
a.11-1	1461	753	0.53	245.4	0.02	0.07565 ± 0.43	2.0391 ± 1.45	0.1955 ± 1.39	0.956	0.0575 ± 1.48	1151 ± 15	1086 ± 9	-6
a.11-2	1036	521	0.52	169.5	0.01	0.07593 ± 0.51	1.9936 ± 1.48	0.1904 ± 1.39	0.939	0.0565 ± 1.67	1124 ± 14	1093 ± 10	-3
a.13-1	1231	668	0.56	198.2	0.01	0.07596 ± 0.46	1.9621 ± 1.46	0.1873 ± 1.38	0.949	0.0569 ± 1.99	1107 ± 14	1094 ± 9	-1
GRT3, zircon													
g.1-1	217	77	0.37	37.3	0.68	0.08133 ± 4.30	2.2308 ± 4.57	0.1989 ± 1.53	0.335	0.0603 ± 8.28	1170 ± 16	1229 ± 84	5
g.2-1a	1284	100	0.08	199.0	0.03	0.07621 ± 0.40	1.8954 ± 1.42	0.1804 ± 1.36	0.959	0.0514 ± 2.83	1069 ± 13	1101 ± 8	3
g.2-1b	1284	100	0.08	199.0	0.00	0.07650 ± 0.38	1.9032 ± 1.42	0.1804 ± 1.36	0.963	0.0531 ± 2.60	1069 ± 13	1108 ± 8	4
g.2-2	76	23	0.31	11.6	0.26	0.07748 ± 1.80	1.8978 ± 2.52	0.1776 ± 1.76	0.700	0.0553 ± 3.45	1054 ± 17	1134 ± 36	7
GRT3, titanite													
d.4-1	226	203	0.93	33.0	0.18	0.07726 ± 2.22	1.8064 ± 2.48	0.1696 ± 1.11	0.449	0.1179 ± 1.82	1010 ± 10	1128 ± 44	10
d.4-2	180	160	0.92	27.5	-0.01	0.07698 ± 0.72	1.8837 ± 1.31	0.1775 ± 1.10	0.835	0.1152 ± 1.27	1053 ± 11	1121 ± 14	6
d.4-3	210	195	0.96	32.4	0.92	0.07621 ± 1.84	1.8748 ± 2.15	0.1784 ± 1.10	0.513	0.0754 ± 1.86	1058 ± 11	1101 ± 37	4

TABLE I (continuation)

spot	U ppm	Th ppm	Th U	$^{206}\text{Pb}/^{206}\text{Pb}$	$4f^{206}\%$	Ratios			Ages			Disc. %	
						$^{207}\text{Pb}/^{235}\text{U}$	$^{206}\text{Pb}/^{238}\text{U}$	err. corr.	$^{208}\text{Pb}/^{232}\text{Th}$	$^{206}\text{Pb}/^{238}\text{U}$	$^{207}\text{Pb}/^{206}\text{Pb}$		
GRT3, titanite													
d.5-1	203	192	0.98	31.3	0.45	0.07511 ± 2.30	1.8508 ± 2.55	0.1787 ± 1.11	0.434	0.1208 ± 1.96	1060 ± 11	1072 ± 46	1
d.5-2	215	197	0.95	32.7	0.25	0.07573 ± 1.37	1.8468 ± 1.76	0.1769 ± 1.10	0.627	0.1049 ± 1.33	1050 ± 11	1088 ± 27	4
d.5-3	184	178	1.00	28.3	0.69	0.07537 ± 2.21	1.8468 ± 2.48	0.1777 ± 1.13	0.454	0.0802 ± 1.63	1054 ± 11	1078 ± 44	2
GRT4, zircon													
e.2-1	129	99	0.79	19.2	0.12	0.07679 ± 1.50	1.8281 ± 3.09	0.1727 ± 2.70	0.874	0.0529 ± 3.04	1027 ± 26	1116 ± 30	8
e.3-1	394	146	0.38	62.2	0.13	0.07595 ± 1.04	1.9228 ± 1.80	0.1836 ± 1.47	0.816	0.0539 ± 2.03	1087 ± 15	1094 ± 21	1
e.3-2	59	30	0.54	9.4	0.95	0.07531 ± 6.50	1.9145 ± 6.79	0.1844 ± 1.97	0.291	0.0526 ± 8.40	1091 ± 20	1077 ± 130	-1.3
e.4-1	527	177	0.35	72.8	0.08	0.07656 ± 0.94	1.6958 ± 1.38	0.1607 ± 1.01	0.733	0.0484 ± 1.48	960 ± 9	1110 ± 19	13
e.4-2	155	76	0.51	23.9	0.02	0.07549 ± 0.90	1.8634 ± 1.41	0.1790 ± 1.09	0.770	0.0545 ± 1.52	1062 ± 11	1082 ± 18	2
e.5-1	297	87	0.30	46.4	0.02	0.07673 ± 1.07	1.9196 ± 1.84	0.1814 ± 1.49	0.812	0.0520 ± 2.48	1075 ± 15	1114 ± 21	4
e.6-1	468	163	0.36	73.7	0.03	0.07653 ± 0.64	1.9355 ± 1.57	0.1834 ± 1.43	0.913	0.0544 ± 1.70	1086 ± 14	1109 ± 13	2
e.7-1	299	298	1.03	45.8	0.13	0.07640 ± 1.07	1.8766 ± 1.82	0.1782 ± 1.47	0.809	0.0530 ± 1.75	1057 ± 14	1105 ± 21	4
e.7-2	37	21	0.60	5.8	0.00	0.07560 ± 2.05	1.9262 ± 2.70	0.1848 ± 1.76	0.651	0.0571 ± 2.83	1093 ± 18	1085 ± 41	-1
GRT5, zircon													
h.1-1	305	57	0.19	48.7	-0.01	0.07661 ± 0.74	1.9625 ± 1.20	0.1858 ± 0.95	0.788	0.0574 ± 1.65	1098 ± 10	1111 ± 15	1
h.2-1	552	89	0.17	89.3	0.01	0.07594 ± 0.79	1.9708 ± 1.26	0.1882 ± 0.98	0.781	0.0544 ± 1.69	1112 ± 10	1093 ± 16	-2
h.2-2	277	70	0.26	44.1	-0.02	0.07645 ± 0.78	1.9568 ± 1.26	0.1856 ± 0.98	0.781	0.0550 ± 1.61	1098 ± 10	1107 ± 16	1
GRT5, rutile													
h.3-3	35	0.893	0.026	4.6	0.39	0.06964 ± 1.4	1.487 ± 3.7	0.1523 ± 3.3	0.92	0.0430 ± 1.29	912 ± 28	953 ± 38	4
h.3-4	35	0.01	0.003	4.8	0.47	0.07388 ± 1.5	1.557 ± 4.0	0.1570 ± 3.0	0.90	0.0513 ± 1.94	942 ± 27	983 ± 52	5
h.6-1	91	0.003	0.003	12.6	0.70	0.07106 ± 2.2	1.588 ± 3.8	0.1610 ± 3.1	0.87	0.0490 ± 1.96	962 ± 39	973 ± 45	1
h.6-2	70.8	0.012	0.001	10.16	0.18	0.07169 ± 2.0	1.604 ± 3.7	0.1667 ± 2.7	0.80	0.0535 ± 1.38	996 ± 25	922 ± 52	-8
h.7-1	85	0.056	0.006	11.8	0.26	0.07114 ± 1.0	1.536 ± 3.7	0.1612 ± 3.2	0.96	0.0540 ± 1.66	966 ± 29	901 ± 35	-7
h.7-2	82	0.056	0.007	11.3	0.43	0.07111 ± 1.0	1.521 ± 3.5	0.1596 ± 3.0	0.95	0.0514 ± 1.35	956 ± 27	903 ± 38	-6
GR752, zircon													
i.2-1	574	131	0.24	134.2	0.15	0.1014 ± 0.7	3.80 ± 1.3	0.2715 ± 1.1	0.84	0.0763 ± 1.6	1552 ± 15	1650 ± 13	6
i.3-1	268	134	0.52	78.9	0.03	0.1203 ± 0.8	5.69 ± 1.4	0.3428 ± 1.2	0.84	0.0976 ± 1.6	1901 ± 21	1960 ± 14	3
i.1-1	605	29	0.05	160.8	0.14	0.1079 ± 0.7	4.59 ± 1.3	0.3086 ± 1.1	0.84	0.1274 ± 2.8	1732 ± 16	1764 ± 13	2
i.1-2	236	98	0.43	62.2	0.10	0.1069 ± 1.0	4.51 ± 1.6	0.3059 ± 1.2	0.79	0.0867 ± 2.3	1724 ± 20	1747 ± 18	2
i.14-1	185	52	0.29	40.7	0.21	0.0883 ± 1.5	3.11 ± 2.3	0.2550 ± 1.8	0.76	0.0738 ± 2.7	1468 ± 24	1390 ± 29	-5
i.14-2	198	218	1.14	62.9	0.23	0.1249 ± 1.1	6.35 ± 1.7	0.3691 ± 1.3	0.77	0.1020 ± 1.6	2039 ± 26	2027 ± 19	0

TABLE I (continuation)

spot	U ppm	Th ppm	Th U	^{206}Pb ppm	$4f^{206}\text{Pb}$ %	Ratios			Ages			Disc. %	
						$^{207}\text{Pb}/^{206}\text{Pb}$	$^{207}\text{Pb}/^{235}\text{U}$	$^{206}\text{Pb}/^{238}\text{U}$	$^{208}\text{Pb}/^{232}\text{Th}$	$^{206}\text{Pb}/^{238}\text{U}$	$^{206}\text{Pb}/^{207}\text{Pb}$		
GR752, zircon													
i.2-2	508	155	0.31	99.8	0.09	0.0915 ± 1.0	2.88 ± 1.2	0.2285 ± 0.7	0.55	0.0673 ± 1.4	1328 ± 8	1457 ± 19	9
i.2-3	319	1	0.00	71.5	0.16	0.0926 ± 0.8	3.33 ± 1.1	0.2608 ± 0.7	0.64	0.1892 ± 7.3	1495 ± 9	1480 ± 16	-1
i.2-4	355	341	0.99	89.5	0.03	0.1095 ± 0.6	4.42 ± 1.0	0.2930 ± 0.7	0.75	0.0849 ± 1.7	1658 ± 13	1791 ± 12	7
i.7-1	580	201	0.36	112.4	0.01	0.0911 ± 0.7	2.83 ± 1.0	0.2256 ± 0.7	0.66	0.0654 ± 1.2	1313 ± 8	1449 ± 14	9
i.7-2	512	254	0.51	124.5	0.04	0.1093 ± 0.9	4.26 ± 1.1	0.2828 ± 0.6	0.57	0.0826 ± 1.6	1606 ± 9	1788 ± 16	10
i.8-1	165	107	0.67	41.5	0.00	0.1072 ± 0.9	4.33 ± 1.3	0.2929 ± 1.0	0.72	0.0789 ± 3.5	1666 ± 16	1753 ± 17	6
i.8-2	627	180	0.30	120.8	0.04	0.0915 ± 0.7	2.83 ± 0.9	0.2243 ± 0.7	0.69	0.0715 ± 1.7	1302 ± 8	1457 ± 13	10
i.8-3	250	67	0.28	51.1	0.16	0.0940 ± 1.2	3.09 ± 1.5	0.2387 ± 0.9	0.58	0.0735 ± 1.9	1376 ± 11	1508 ± 23	8
i.9-1	329	144	0.45	67.5	0.01	0.0912 ± 0.8	3.00 ± 1.1	0.2386 ± 0.8	0.70	0.0708 ± 1.3	1379 ± 10	1451 ± 15	5
i.3-2	35	18	0.53	10.8	0.19	0.1208 ± 1.8	6.05 ± 2.8	0.3629 ± 2.1	0.76	0.0984 ± 3.2	2006 ± 39	1968 ± 33	-1
i.6-1	250	96	0.40	66.3	0.04	0.1083 ± 0.7	4.60 ± 1.5	0.3081 ± 1.3	0.88	0.0890 ± 1.7	1733 ± 21	1772 ± 13	2
i.5-1	175	159	0.94	56.0	0.04	0.1224 ± 0.7	6.30 ± 1.5	0.3736 ± 1.4	0.89	0.1076 ± 2.1	2045 ± 27	1991 ± 12	-3
i.15-1	249	189	0.78	69.5	0.13	0.1081 ± 0.7	4.83 ± 1.5	0.3244 ± 1.3	0.87	0.0929 ± 1.7	1815 ± 23	1767 ± 14	-2
i.13-1	405	174	0.44	90.1	0.05	0.0909 ± 0.6	3.24 ± 1.4	0.2587 ± 1.2	0.89	0.0733 ± 1.5	1487 ± 17	1444 ± 12	-3
i.13-2	240	71	0.31	46.5	0.21	0.0919 ± 1.2	2.85 ± 1.8	0.2252 ± 1.3	0.74	0.0766 ± 1.8	1305 ± 17	1465 ± 23	11
i.10-1	455	217	0.49	101.4	0.03	0.0910 ± 0.6	3.26 ± 1.4	0.2594 ± 1.2	0.90	0.0713 ± 1.4	1493 ± 17	1448 ± 11	-3
NM165, zircon													
a.2-1	389	92	0.24	80.3	-0.01	0.08647 ± 0.74	2.8678 ± 1.65	0.2405 ± 1.48	0.895	0.0660 ± 2.14	1390 ± 19	1349 ± 14	-3
a.3-1	745	387	0.54	168.0	0.29	0.09420 ± 0.50	3.3994 ± 1.10	0.2616 ± 0.99	0.894	0.0731 ± 1.27	1498 ± 13	1513 ± 9	1
a.3-2	781	336	0.44	178.3	0.13	0.09457 ± 0.56	3.4592 ± 1.13	0.2652 ± 0.99	0.871	0.0744 ± 1.22	1516 ± 13	1520 ± 11	0
a.5-1	930	550	0.61	214.6	0.01	0.09495 ± 0.45	3.5162 ± 1.09	0.2686 ± 0.99	0.911	0.0765 ± 1.10	1534 ± 14	1527 ± 8	0
a.8-1	1324	41	0.03	258.2	0.00	0.08673 ± 0.44	2.7150 ± 1.54	0.2270 ± 1.48	0.959	0.0621 ± 3.65	1319 ± 18	1354 ± 8	3
a.11-1	1171	11	0.01	232.3	0.03	0.08615 ± 0.34	2.7408 ± 1.02	0.2307 ± 0.97	0.944	0.0576 ± 7.99	1338 ± 12	1342 ± 7	0
a.12-1	993	102	0.11	196.2	0.05	0.08569 ± 0.40	2.7147 ± 1.05	0.2297 ± 0.97	0.926	0.0633 ± 2.04	1333 ± 12	1331 ± 8	0
a.13-1	474	162	0.35	107.9	-0.01	0.09389 ± 0.63	3.4342 ± 1.14	0.2653 ± 0.95	0.834	0.0782 ± 1.28	1517 ± 13	1506 ± 12	-1
a.13-2	1218	185	0.16	262.7	0.04	0.09051 ± 0.31	3.1307 ± 1.01	0.2509 ± 0.96	0.952	0.0756 ± 1.25	1443 ± 12	1436 ± 6	0
a.14-1	1088	38	0.04	213.9	0.01	0.08637 ± 0.42	2.7245 ± 1.07	0.2288 ± 0.98	0.919	0.0659 ± 2.86	1328 ± 12	1346 ± 8	1
a.16-1	90	33	0.38	17.8	0.00	0.08583 ± 1.41	2.7313 ± 2.07	0.2308 ± 1.52	0.732	0.0682 ± 2.51	1339 ± 18	1334 ± 27	0
GR768, zircon													
d-iv-2	1817	188	0.10	384	0.003	0.0979 ± 0.77	3.4721 ± 0.92	0.2573 ± 0.50	0.55	1476 ± 7	1476 ± 7	1584 ± 12	7
d-iv-3	352	98	0.28	80	0.001	0.0983 ± 1.34	3.3567 ± 1.47	0.2477 ± 0.60	0.41	1427 ± 8	1427 ± 8	1592 ± 21	10
d-iv-4	1576	129	0.08	365	0.017	0.0986 ± 1.67	3.2932 ± 1.80	0.2421 ± 0.67	0.37	1398 ± 9	1398 ± 9	1599 ± 27	13

TABLE I (continuation)

spot	U ppm	Th ppm	Th U	$^{206}\text{Pb}/^{206}\text{Pb}$	$4f^{206}\text{Pb}/^{206}\text{Pb}$ %	Ratios				Ages		Disc. %	
						$^{207}\text{Pb}/^{206}\text{Pb}$	$^{207}\text{Pb}/^{235}\text{U}$	$^{206}\text{Pb}/^{238}\text{U}$	err. corr.	$^{208}\text{Pb}/^{232}\text{Th}$	$^{206}\text{Pb}/^{238}\text{U}$		$^{206}\text{Pb}/^{207}\text{Pb}$
GR768, zircon													
d-iv-5	2197	217	0.10	491	0.002	0.0958 ± 1.75	3.3416 ± 1.82	0.2530 ± 0.52	0.29	1454 ± 8	1454 ± 8	1544 ± 27	6
d-iv-7	1556	137	0.09	337	0.001	0.0944 ± 0.87	3.3152 ± 1.07	0.2548 ± 0.61	0.57	1463 ± 9	1463 ± 9	1516 ± 13	3
e-v-2	1870	348	0.19	350	0.002	0.0941 ± 1.33	3.2910 ± 2.03	0.2536 ± 1.53	0.76	1457 ± 22	1457 ± 22	1510 ± 20	4
e-v-3	148	38	0.26	24	0.001	0.0737 ± 1.26	1.7552 ± 1.31	0.1727 ± 0.34	0.26	1027 ± 3	1027 ± 3	1033 ± 13	1
e-v-4	26	4	0.17	4	0.002	0.0731 ± 2.80	1.7090 ± 2.99	0.1696 ± 1.04	0.35	1010 ± 10	1010 ± 10	1017 ± 29	1
e-v-5	108	28	0.26	17	0.003	0.0735 ± 1.40	1.6898 ± 1.49	0.1667 ± 0.52	0.35	994 ± 5	994 ± 5	1028 ± 14	3
e-v-6	84	16	0.19	13	0.003	0.0731 ± 1.49	1.6941 ± 1.63	0.1680 ± 0.65	0.40	1001 ± 6	1001 ± 6	1018 ± 15	2
GR66, zircon													
d.54-1	447	100	0.223	0.0001	0.085	0.0535 ± 13	2.277 ± 42	0.2039 ± 29	0.90	0.0489 ± 14	1196 ± 16	1222 ± 20	2
d.56-1	422	185	0.438	0.0001	0.093	0.1300 ± 17	2.453 ± 56	0.2165 ± 42	0.74	0.0643 ± 15	1263 ± 22	1250 ± 19	-1
d.57-1	255	39	0.153	0.000	0.002	0.0460 ± 20	2.729 ± 70	0.2315 ± 46	0.84	0.0696 ± 33	1342 ± 24	1327 ± 27	-1
d.58-1	64	34	0.535	0.000	0.062	0.1583 ± 64	2.326 ± 102	0.2128 ± 48	0.87	0.0630 ± 30	1244 ± 26	1179 ± 69	-5
d.59-1	466	98	0.210	0.000	0.070	0.0618 ± 12	2.916 ± 64	0.2375 ± 46	0.76	0.0699 ± 20	1373 ± 24	1405 ± 16	2
d.60-1	845	162	0.192	0.000	0.001	0.0576 ± 6	3.013 ± 61	0.2430 ± 46	0.70	0.0729 ± 16	1402 ± 24	1424 ± 10	2
d.61-1	140	90	0.646	0.000	0.016	0.1787 ± 23	6.288 ± 150	0.3666 ± 75	0.86	0.1015 ± 25	2013 ± 35	2020 ± 17	0
d.62-1	61	21	0.350	0.000	0.000	0.1066 ± 32	2.200 ± 74	0.2002 ± 46	0.63	0.0609 ± 24	1177 ± 25	1189 ± 43	1
d.63-1	73	31	0.426	0.0003	0.405	0.1184 ± 75	2.105 ± 109	0.2001 ± 45	0.75	0.0556 ± 38	1176 ± 24	1103 ± 88	-7
d.64-1	103	62	0.598	0.0001	0.238	0.1712 ± 44	2.532 ± 84	0.2197 ± 47	0.66	0.0629 ± 22	1280 ± 25	1283 ± 44	0
d.65-1	71	43	0.604	0.000	0.075	0.1718 ± 46	3.619 ± 119	0.2797 ± 61	0.57	0.0796 ± 28	1590 ± 31	1505 ± 41	-6
d.67-1	113	37	0.326	0.0002	0.378	0.0936 ± 63	2.644 ± 111	0.2265 ± 48	0.72	0.0650 ± 46	1316 ± 25	1308 ± 66	-1
d.69-1	274	207	0.756	0.0001	0.147	0.1019 ± 18	4.265 ± 97	0.2943 ± 58	0.69	0.0396 ± 11	1663 ± 29	1716 ± 17	3
d.70-1	672	332	0.495	0.000	0.024	0.1438 ± 12	2.217 ± 47	0.1995 ± 38	0.58	0.0580 ± 12	1173 ± 21	1211 ± 14	3
d.71-1	99	56	0.571	0.000	0.069	0.1675 ± 44	2.292 ± 78	0.2151 ± 46	0.76	0.0631 ± 22	1256 ± 24	1122 ± 48	-11
d.73-1	226	65	0.285	0.000	0.059	0.0840 ± 23	2.406 ± 63	0.2163 ± 43	0.79	0.0637 ± 22	1262 ± 23	1213 ± 29	-4
d.75-1	224	95	0.389	0.0001	0.000	0.1020 ± 16	1.991 ± 48	0.1817 ± 36	0.88	0.0476 ± 12	1076 ± 20	1184 ± 22	9
d60.11	179	37	0.209	0.0004	0.616	0.0559 ± 45	2.929 ± 90	0.2429 ± 39	0.55	0.0649 ± 53	1402 ± 20	1370 ± 47	-2
d.61.2	171	75	0.440	0.000	0.008	0.1318 ± 22	6.746 ± 131	0.3781 ± 59	0.84	0.1133 ± 27	2068 ± 28	2090 ± 17	1
d65.11	76	52	0.683	0.0001	0.108	0.1959 ± 53	3.687 ± 117	0.2709 ± 52	0.89	0.0777 ± 27	1546 ± 26	1599 ± 43	3
d69.11	291	123	0.422	0.000	0.050	0.1130 ± 18	4.706 ± 87	0.3191 ± 48	0.74	0.0854 ± 20	1785 ± 23	1748 ± 16	-2
d.74-1	288	93	0.321	0.000	0.068	0.0933 ± 24	2.281 ± 50	0.1940 ± 29	0.90	0.0564 ± 17	1143 ± 16	1322 ± 27	13
d.76-1	237	95	0.399	0.0001	0.093	0.1211 ± 22	2.648 ± 57	0.2244 ± 35	0.84	0.0681 ± 17	1305 ± 18	1329 ± 25	2
d.77-1	215	185	0.858	0.000	0.047	0.2491 ± 26	5.515 ± 105	0.3432 ± 53	0.57	0.0996 ± 100	1902 ± 25	1904 ± 17	0

TABLE I (continuation)

spot	Isotopic ratios .			Ages				f206	Dis %	²³² Th / ²³⁸ U
	²⁰⁷ Pb . / ²³⁵ U	²⁰⁶ Pb . / ²³⁸ U	Rho	²⁰⁷ Pb . / ²⁰⁶ Pb	²⁰⁶ Pb . / ²³⁸ U	²⁰⁷ Pb . / ²³⁵ U	²⁰⁷ Pb . / ²⁰⁶ Pb			
NM182, zircon										
cv-01	2.53454 ± 1.41	0.21420 ± 0.52	0.37	0.08582 ± 1.31	1251 ± 7	1282 ± 18	1334 ± 18	0.0004	6	0.38
cv-02	3.17845 ± 1.40	0.24659 ± 0.45	0.32	0.09348 ± 1.32	1421 ± 6	1452 ± 20	1498 ± 20	0.0001	5	0.43
cv-04	5.97356 ± 2.94	0.35067 ± 0.32	0.11	0.12355 ± 2.93	1938 ± 6	1972 ± 58	2008 ± 59	0.0000	4	0.50
cv-39	3.46876 ± 1.21	0.26243 ± 0.64	0.53	0.09586 ± 1.02	1502 ± 10	1520 ± 18	1545 ± 16	0.0006	3	0.20
cv-40	3.40291 ± 0.87	0.26381 ± 0.46	0.53	0.09355 ± 0.74	1509 ± 7	1505 ± 13	1499 ± 11	0.0001	-1	0.35
cv-42	3.44029 ± 0.94	0.26608 ± 0.56	0.59	0.09377 ± 0.75	1521 ± 8	1514 ± 14	1504 ± 11	0.0001	-1	0.50
cv-43	4.03463 ± 1.16	0.29477 ± 0.81	0.70	0.09927 ± 0.83	1665 ± 13	1641 ± 19	1610 ± 13	0.0002	-3	0.69
cv-48	3.63067 ± 1.76	0.27049 ± 0.77	0.44	0.09735 ± 1.59	1543 ± 12	1556 ± 27	1574 ± 25	0.0006	2	0.72
cv-51	3.49946 ± 1.24	0.27216 ± 0.95	0.77	0.09325 ± 0.79	1552 ± 15	1527 ± 19	1493 ± 12	0.0001	-4	0.34
cv-56	3.47862 ± 1.01	0.26923 ± 0.54	0.53	0.09371 ± 0.86	1537 ± 8	1522 ± 15	1502 ± 13	0.0008	-2	0.34
GR336, zircon										
b-ii-1	3.39949 ± 1.13	0.26236 ± 0.76	0.68	0.09397 ± 0.83	1502 ± 11	1504 ± 17	1508 ± 12	0.0003	0	0.30
b-ii-3	3.48677 ± 0.95	0.26728 ± 0.72	0.76	0.09462 ± 0.62	1527 ± 11	1524 ± 14	1520 ± 9	0.0003	0	0.25
b-ii-5	3.55896 ± 1.91	0.26826 ± 1.28	0.67	0.09622 ± 1.42	1532 ± 20	1540 ± 29	1552 ± 22	0.0014	1	0.51
b-ii-6	2.84822 ± 1.15	0.23945 ± 0.78	0.68	0.08627 ± 0.84	1384 ± 11	1368 ± 16	1344 ± 11	0.0001	-3	0.32
b-ii-7	1.66614 ± 0.98	0.16179 ± 0.63	0.64	0.07469 ± 0.75	967 ± 6	996 ± 10	1060 ± 8	0.0001	9	0.06
b-ii-2	2.76130 ± 1.34	0.23151 ± 0.77	0.58	0.08650 ± 1.09	1342 ± 10	1345 ± 18	1349 ± 15	0.0005	1	0.33
b-iii-3	2.81367 ± 1.44	0.23971 ± 1.11	0.77	0.08513 ± 0.92	1385 ± 15	1359 ± 20	1319 ± 12	0.0003	-5	0.78
b-iii-6	2.78383 ± 1.32	0.23658 ± 0.84	0.63	0.08534 ± 1.03	1369 ± 11	1351 ± 18	1323 ± 14	0.0003	-3	0.28
biii7a	3.52090 ± 1.05	0.27001 ± 0.77	0.73	0.09457 ± 0.72	1541 ± 12	1532 ± 16	1520 ± 11	0.0001	-1	0.35
biii7b	1.68423 ± 1.84	0.16244 ± 1.73	0.94	0.07520 ± 0.64	970 ± 17	1003 ± 18	1074 ± 7	0.0001	10	0.04
biii11	2.62565 ± 1.36	0.22332 ± 0.91	0.67	0.08527 ± 1.02	1299 ± 12	1308 ± 18	1322 ± 13	0.0001	2	0.44
b-iv-1	2.64334 ± 1.09	0.22359 ± 0.65	0.60	0.08574 ± 0.88	1301 ± 8	1313 ± 14	1332 ± 12	0.0001	2	0.25
b-iv-3	3.03591 ± 1.18	0.24366 ± 1.00	0.85	0.09037 ± 0.62	1406 ± 14	1417 ± 17	1433 ± 9	0.0003	2	0.27
b-iv-5	2.78782 ± 1.12	0.23325 ± 0.92	0.83	0.08669 ± 0.63	1351 ± 12	1352 ± 15	1354 ± 9	0.0001	0	0.28
b-iv-7	2.81440 ± 1.15	0.23754 ± 0.79	0.69	0.08593 ± 0.83	1374 ± 11	1359 ± 16	1337 ± 11	0.0003	-3	0.28
GR345, zircon										
a-i-3	1.71977 ± 1.33	0.16868 ± 0.77	0.58	0.07394 ± 1.08	1005 ± 8	1016 ± 13	1040 ± 11	0.0061	3	0.74
a-i-5	1.76550 ± 1.23	0.17646 ± 0.49	0.40	0.07257 ± 1.13	1048 ± 5	1033 ± 13	1002 ± 11	0.0001	-5	0.35
a-i-7	1.76088 ± 1.26	0.17620 ± 0.74	0.59	0.07248 ± 1.02	1046 ± 8	1031 ± 13	999 ± 10	0.0008	-5	0.62
a-i-12	1.53957 ± 1.75	0.15376 ± 0.71	0.41	0.07262 ± 1.60	922 ± 7	946 ± 17	1003 ± 16	0.0003	8	0.46
a-i-13	1.76515 ± 1.42	0.17634 ± 0.99	0.70	0.07260 ± 1.02	1047 ± 10	1033 ± 15	1003 ± 10	0.0002	-4	0.75
a-i-16	1.72609 ± 1.59	0.17265 ± 0.83	0.52	0.07251 ± 1.36	1027 ± 8	1018 ± 16	1000 ± 14	0.0009	-3	0.50
a-i-25	1.83594 ± 1.54	0.17989 ± 1.12	0.73	0.07402 ± 1.06	1066 ± 12	1058 ± 16	1042 ± 11	0.0144	-2	0.51
a-i-28	1.74057 ± 1.57	0.17504 ± 0.93	0.59	0.07212 ± 1.27	1040 ± 10	1024 ± 16	989 ± 13	0.0001	-5	0.50
a-i-32	1.71426 ± 1.53	0.17251 ± 0.74	0.48	0.07207 ± 1.34	1026 ± 8	1014 ± 16	988 ± 13	0.0001	-4	0.99
a-i-34	1.75302 ± 1.37	0.17414 ± 0.75	0.55	0.07301 ± 1.14	1035 ± 8	1028 ± 14	1014 ± 12	0.0004	-2	0.75

Notes: Isotopic ratios errors in %; All Pb in ratios are radiogenic component. Most are corrected for ²⁰⁴Pb and some for ²⁰⁸Pb (metamorphic, Th-poor grains or rims). disc. = discordance, as 100 - 100{t[²⁰⁶Pb/²³⁸U]/t[²⁰⁷Pb/²⁰⁶Pb]}; f206 = (common ²⁰⁶Pb) / (total measured ²⁰⁶Pb) based on measured ²⁰⁴Pb. Uncertainties are 1σ.

has an elongated granoblastic texture. However, later hydrous retrogression is widespread along transcurrent shear zone, and the orthopyroxene shows various degrees of alteration to amphibolite. The analyzed samples show minimal retrogression.

Sample GRT-01 (S 11°57'49"/W 61°44'58") is a dark gray, coarse-grained mafic granulite with granoblastic texture. It is fine-grained within mylonitic zones and displays several millimeter-to-centimeter-sized veinlets of amphibole that are randomly oriented and associated with sulfides, resulting from retrogressive amphibolite facies conditions. Their mylonitic foliation trends N60°W; 30°NE, with mineral lineation at 35°; N30°W.

Zircon grains from this sample are generally of 80 to 120 μm in size with partly rounded or broken terminations and are generally structureless or patchily zoned. Zircon grains reveal internal features characteristic of the solid state recrystallization, even on grains from rocks in which igneous textures are preserved. This alteration process is pervasive through the mafic rocks of the belt.

The U content is highly variable (706 to 2,254 ppm) as is the Th content which ranges between 254 and 2,266 ppm (Table I). From a population of 5 analyzed crystals, 4 crystals yielded suitable isotopic response, indicating the upper intercept age of 1114 ± 5 Ma, with an MSWD of 1.09 (Fig. 4a), which is interpreted as peak metamorphism age. This age is in agreement with the age measured in the anatectic melt (1110 ± 8 Ma) of the northern domain of the Nova Brasilândia belt (Rizzotto 1999). However, there is one zircon (grain h.1-3) with discordant ages near 1014 Ma, which may indicate Pb loss due to the superimposition of a later metamorphic event.

Sample GRT-02 (S 11°57'04"/W 61°45'01") is a banded leucocratic trondhjemite, fine- to medium-grained, and forms subconcordant lenses (ca. 300 m long and 50 m thick) intercalated with mafic granulite. It consists mainly of quartz and variably saussuritized plagioclase, with $\leq 10\%$

interstitial mafic minerals (mostly amphibole and clinopyroxene, with titanite, magnetite and traces of zircon). Titanite crystals contain inclusions of biotite, plagioclase, quartz and epidote. Titanite crystals develop along the foliation planes characterizing synkinematic crystals. Much of the matrix titanite is aligned and concentrated in amphibole-rich bands that anastomose around granoblastic plagioclase/quartz-rich domains.

Zircon and titanite crystals were dated. Zircons grains are light tan to colorless, clear and show a varied morphology with sizes ranging from 100 μm to 350 μm . Elongated prisms characterize the most frequent zircon type, but ovoid habits with rounded terminations are also observed. The zircon crystals reveal the most varied internal features, with a disruption of concentric oscillatory zonation resulting in zoning irregular patches and ghost zoning. Some grains have fractured cores and homogeneous rims.

The U content of zircon varies from 509 to 1,461 ppm (Table I). Thirteen analyses in ten zircon grains were obtained from this sample. Eight of these analyses yielded a weighted mean $^{207}\text{Pb}/^{206}\text{Pb}$ age of 1089 ± 9 Ma (MSWD = 1.5 - Fig. 4b), which is interpreted as the metamorphic age of the trondhjemite. The age of metamorphism is slightly younger than that of sample GRT-01, which may be due to Pb loss resulting from alteration processes such as metamictization.

Light tan to dark titanite grains are subhedral to rounded and unzoned, and generally, crystals are fragmented with widths that vary from 120 to 200 μm . The U contents of the grains are generally lower than 300 ppm, and the Th content varies between 180 and 260 ppm. Regardless, the Th/U ratios are high, and their occurrence within metamorphic bands indicates that these grains grew in a metamorphic environment.

The five analyzed grains plot on concordia (Fig. 4b), indicating a concordant age of 986 ± 6 Ma. The large difference between the titanite age

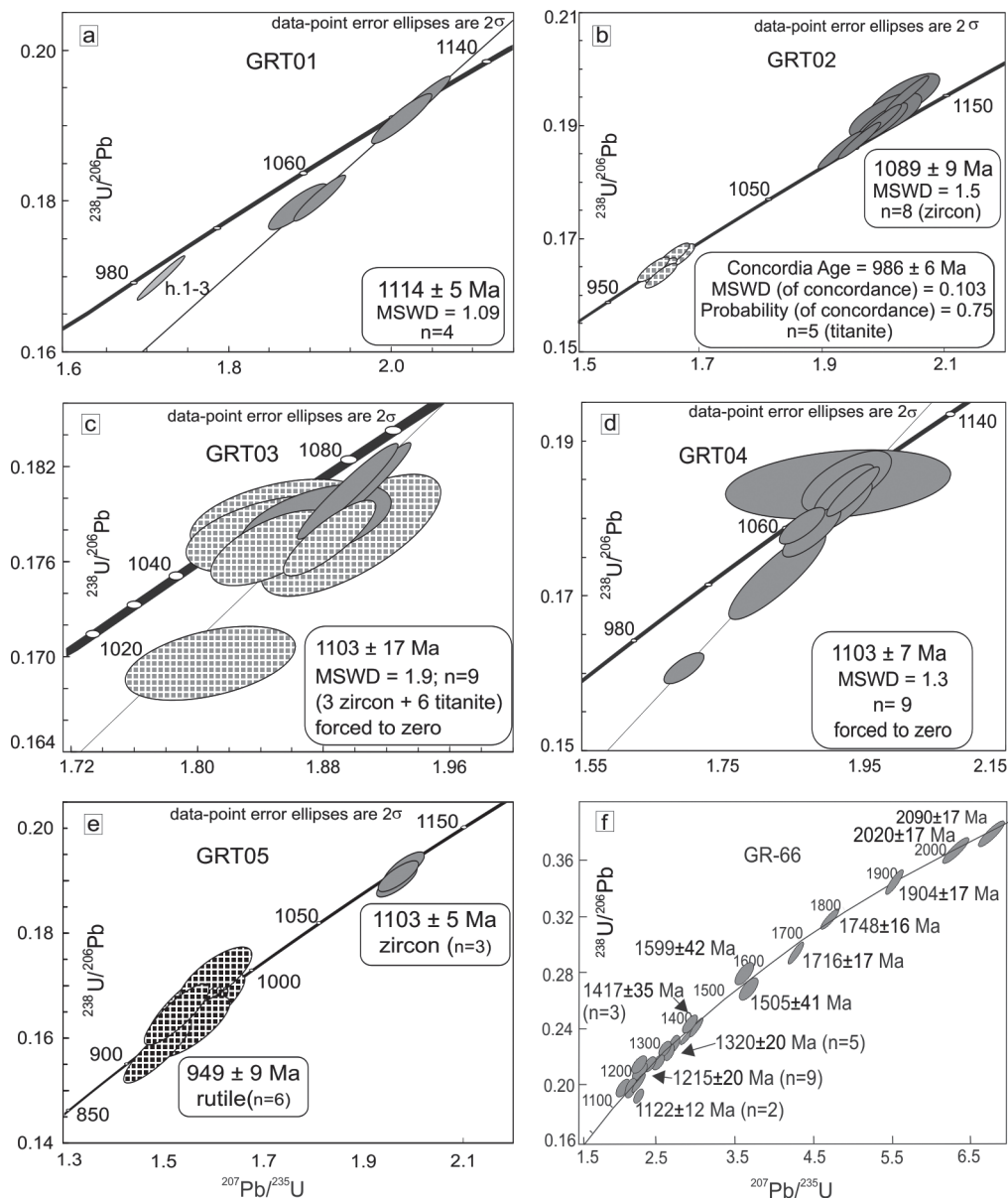


Figure 4 - Concordia diagrams of dated samples of the Nova Brasilândia belt.

and zircon age suggests that there may have been a deformation phase at ~ 1.0 Ga. Another possibility is that titanite has a lower closure temperature than zircon, and thus, this age would represent the cooling after peak metamorphism. In this same area, Tohver et al. (2004) dated monazite crystals from a mylonitic gneiss, near GRT-01, which yielded the age of 993 ± 11 Ma. These authors argue that monazite grew below its closure temperature through dynamic

recrystallization. This interpretation is in agreement with the previous suggestion of a later phase deformation at ca. 1.0 Ga.

Sample GRT-03 (S $11^{\circ}53'11''$ /W $62^{\circ}29'11''$) is a light gray banded trondhjemite, medium-grained, in decametric lenses intercalated with mafic granulites. The mineralogy and texture of this trondhjemite are similar to sample GRT-02. However, it has higher content of mafic minerals

(amphibole and clinopyroxene) and banding defined by the preferred orientation of these minerals. The metamorphic fabrics range in intensity from a well-developed foliation to gneissic banding in high strain zones. Lineation is characterized by low rake angle, suggestive of a strike-slip component of the motion.

Zircon crystals are slightly elongate and exhibit stubby habit with rounded terminations, which render oval morphologies to some grains. Generally, crystals are fragmented, and their length varies between 100 and 200 μm . Irregular light patches and ghost zoning are also commonly observed. Four zircon grains were analyzed and show high U contents between 217 and 1,284 ppm, although one analysis yielded only 76 ppm. Th contents are mostly high, varying between 23 and 100 ppm. Low Th/U ratios (0.08 – Table I) indicate a correlation with light patches, suggesting that these zones are metamorphic. One zircon grain with a core with irregular concentric oscillatory zoning in the core yielded the oldest age (1229 Ma), either magmatic or inherited. In this scenario, the mafic rocks would have assimilated zircon from an older crust, although rocks of this age not been discovered in the region.

In addition to of zircons, six titanite grains were analyzed. They are generally fragmented with various sizes and shapes, and their lengths vary between 200 and 500 μm . Most of the grains are unzoned or have faint, patchy zoning. Titanite crystals have U contents lower than 210 ppm, varying between 180 and 203 ppm, and have Th near 200 ppm. Combining the data from zircon and titanite, the metamorphic recrystallization in this sample is dated at 1103 ± 17 Ma (Fig. 4c) by nine analyses (3 zircon + 6 titanite).

Fine-grained amphibolite (*Sample GRT-04*; S $11^{\circ}53'23''$ /W $62^{\circ}29'16''$) of upper amphibolite facies with minor intercalations of trondhjemite dominates the central part of the mafic belt of the Nova Brasilândia proto-ocean. This sample is amphibole-rich amphibolite that partially

preserves primary igneous features. In thin section, brown and clear titanite grains occur primarily in amphibole-epidote-rich domains. The plagioclase has lobate grain boundaries, often recrystallized to small subgrains, and shows evidence of strong deformation. The mineralogical and textural features suggest gabbro as the protolith.

Zircon grains this sample are colorless, clear, elongated, prismatic-to-rounded, and many contain numerous cracks. Most grains are structureless or patchily zoned. Irregular light patches occur in various parts of the crystals. The U content of the zircon varies between 37 and 527 ppm, and the Th contents between 21 and 298 ppm. Nine analyses this sample yield an intercept age of 1103 ± 7 Ma (MSWD= 1.3 - Fig. 4d), which is interpreted as the age of peak metamorphism. This age is the same as that obtained from sample GRT-03 described above.

A coarse-grained amphibolite was collected at the westernmost part of the mafic belt of the Nova Brasilândia rift, near the contact with a porphyroblastic milonitic granite. *Sample GRT-05* (S $11^{\circ}57'08''$ /W $63^{\circ}05'57''$) has partially preserved igneous texture and zones with well-developed mylonitic foliation. The mylonitic fabric is defined by a preferred orientation of amphibole or by sub-centimeter-to-centimeter scale titanite-ilmenite veinlets. Their mylonitic foliation shows a W-E structural trend.

Zircon and rutile crystals were analyzed. Zircon crystals are weakly zoned (ghost zoning) and generally 100 to 200 μm in size. They exhibit stubby habit with rounded terminations, which provide oval-to-rounded morphologies to some grains. U contents range from 277 to 552 ppm, and Th contents are more uniform and vary from 57 to 89 ppm. U–Pb isotopic analyses of three zircon grains yield an average $^{207}\text{Pb}/^{206}\text{Pb}$ age of 1103 ± 5 Ma (Fig. 4e), which we interpret to be a metamorphic recrystallization age.

Rutile grains show an external morphology that varies from prismatic to slightly rounded, and

these grains are 100-250 μm long. Six grains were dated and yielded a $^{207}\text{Pb}/^{206}\text{Pb}$ concordia age of 949 ± 9 Ma (Fig. 4e). Taking into account the analytical uncertainty, this age is similar to the titanite age of sample GRT-02. Generally, the U/Pb ages for rutile are younger than the age of peak metamorphism and younger than the U/Pb ages for zircon, titanite and monazite (Mezger et al. 1989). This result suggests that the rutile ages reflect cooling below closure temperatures. However, there is a large difference between the peak metamorphic age recorded by zircon (~ 1100 Ma) and the age of metamorphic rutile (~ 950 Ma). Another possible scenario is that the metamorphic rutile grew below its closure temperature through dynamic recrystallization, promoted by a later phase of deformation at ca. 1.0 Ga. This later phase of deformation is related to the Rio Branco transcurrent shear zone development in low amphibolite facies metamorphic conditions (Rizzotto 1999).

The sample GR-66 (S $11^{\circ}43'33''$ /W $62^{\circ}09'10''$) was previously dated by Rizzotto (1999) and has now been reprocessed. It is a fine-grained banded paragneiss, with well-defined felsic bands composed of quartz and plagioclase and mafic bands of biotite. The mineral assemblage is in metamorphic equilibrium compatible with upper amphibolite facies conditions.

Twenty-four grains from this sample were analyzed (Table I). The largest population (9 grains) yield an age of 1215 ± 20 Ma. Five grains yield older ages of 1320 ± 20 Ma, while a group of three crystals has an age of 1417 ± 35 Ma. A younger population with an age of 1122 ± 12 Ma was also obtained (2 grains). The remaining data yield a spread of ages between 2090 and 1505 Ma (Fig. 4f). These results suggest that the protolith of the Nova Brasilândia paragneiss was derived from a Paleoproterozoic/Mesoproterozoic mixed source. These data further suggest that the sedimentary protolith of the Nova Brasilândia Group was not deposited prior to 1122 Ma, the youngest age obtained from the population of detrital zircon grains.

Therefore, all of the ages reported above are equivalent, within error ranges, and represent the timing of peak metamorphism (~ 1104 Ma) after the late Mesoproterozoic intraplate rifting (1150-1122 Ma).

AGUAPEÍ BELT

The sampled area is the most northwestern region of the belt, near the southwestern boundary of the Parecis Basin (Figs. 1 and 2). This area consists of a repetitive, intercalated sequence of pelitic and quartzose layers, suggesting a turbiditic sequence. In spite of the mid-to-low metamorphic grade of this area (greenschist to lower amphibolite facies), S_0 bedding is locally preserved and still distinguishable in some outcrops. The metasedimentary rocks of the area investigated in this work are structurally and lithologically distinct from the mature, continental sediments of the Aguapeí Group (e.g., Souza and Hildred 1980, Saes 1999). However, we propose that these metasedimentary rocks (sand-silt rhythmites) form part of the basal unit of the Aguapeí Group. Accordingly, this metasedimentary package is very similar structurally and lithologically to the metasedimentary rocks of the Nova Brasilândia Group.

Sample NM-165 (S $13^{\circ}25'39''$ /W $60^{\circ}15'20''$) is an aluminous metapsammite, which is interbedded with metapelitic layers, suggesting a turbiditic environment. It shows pervasive symmetrical folds with mainly quartz, biotite and garnet and fine veinlets of quartz along the axial plane.

Zircon grains from this sample vary in size with partly rounded or have broken terminations, and vary in size from 100-300 μm with an aspect ratio of 3:1 to 2:1. Although magmatic euhedral faces are present in some crystals, rounding is a common feature in many crystals, and all are detrital. The main population ($n=6$) has an age in the 1331-1354 Ma range (Table I and Fig. 5a) interpreted to have been derived from rocks of the Alto Guaporé belt, such as the granitoids of the Igarapé Enganado and Alto Escondido suites (Rizzotto 2010). Four grains

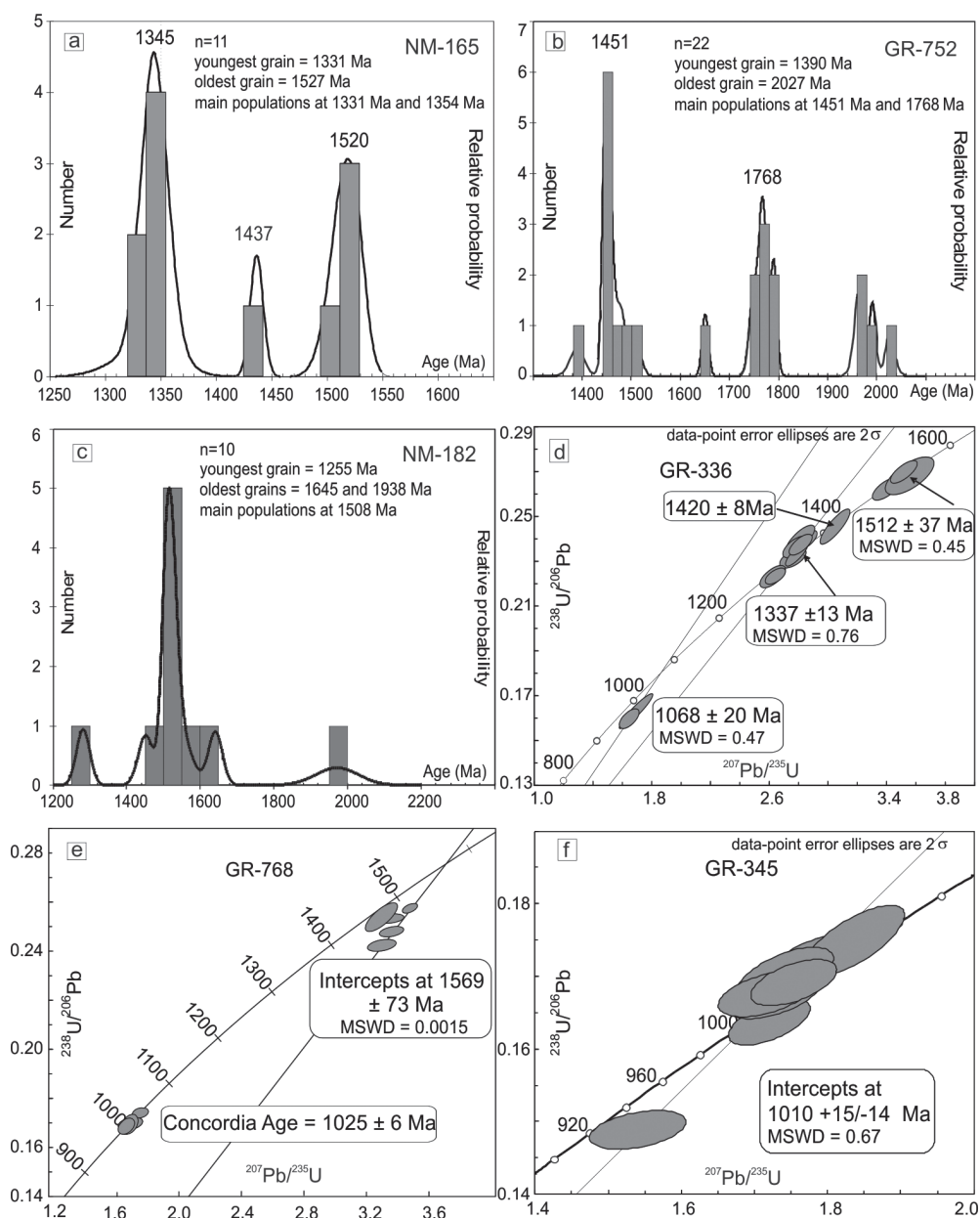


Figure 5 - U-Pb concordia plots and relative probability plots of zircon ages from samples of the Aguapei belt and intrusives rocks.

have ages between 1506 and 1527 Ma, derived from rocks of the Rio Alegre terrane. One grain has an age of 1436 ± 6 Ma and derived from the Trinchera complex. The youngest zircon yielded an age of 1331 ± 8 Ma (Table I), which is interpreted as the maximum sedimentation age of the sedimentary protolith.

Sample GR-752 (S $13^{\circ}25'26''$ /W $60^{\circ}15'40''$) is a banded phyllite interbedded with quartzite layers.

It consists of quartz, biotite and garnet aggregates arranged along a well-developed vertical foliation. Many ptigmatic folds consists mainly of quartz, and the layers are cross-cut by an dense network of quartz veins. Folds axis are consistent with the mineral lineation at 70° ; $S80^{\circ}W$.

Most zircon grains are short prisms with rounded edges and variable shapes and color and

are all detrital. They present two main populations, 1451 and 1768 Ma; the youngest grain has an age of 1390 Ma and the oldest grain 2027 Ma (Fig. 5b). The ages of the most representative zircon populations (1.45 Ga; 1.76 Ga) signify not only the different continental sedimentation sources (Leite and Saes 2003) but also the timing of the adjacent magmatic arcs (Jamari and Santa Helena magmatic arcs), which forms part of the basement in the Rondônia and Mato Grosso states of the southwestern margin of the Amazonian craton.

Sample NM-182 (S 13°26'35"/W 60°09'25") is an impure quartzite (metagreywacke), which presents overturned and subvertical beds with metric to decimetric thickness and oriented banding at N05°E/85°SE, although it is strongly folded in the highest strain zones. This quartzite is interbedded with thin beds of biotite-rich metamudstones (phyllites).

Most zircon grains are small (80-120 µm) with rounded edges, and several types of shapes are present. The results obtained for the 10 analyzed zircon crystals (Table I) show that the quartzite has a wide range of zircon provenances (Fig. 5c). Most of the crystals (n=7) have an age of 1508 Ma and may have been derived from rocks correlated to the Cachoeirinha orogeny (Ruiz 2005). Two grains yield older ages of 1938 and 1645 Ma and are interpreted to have been derived from rocks constituting the basement of the Rondônia-Juruena Province. One grain has a younger age (1255 Ma), which may indicate the maximum depositional age of the protoliths of the Aguapeí belt and most likely was derived from rocks correlated to the San Andres Granite (Santos et al. 2008). This granite occurs in the easternmost part of Bolivia or the provenance is from another granitoid body within the basement of the Rondônia-Juruena Province.

Sample GR336 (S 11°59'02"/W 61°23'58") is a fine-grained dark banded gneiss, which contains abundant elongated quartz nodules, in a strongly foliated matrix of quartz-plagioclase-biotite-ilmenite, including significant areas of leucosome development.

The zircon population is dominated by rounded grains, short prisms, and a very small proportion of the prismatic grains. They present variable shapes and color and are all detrital. The main population (n=8) has an age of 1337 Ma (Fig. 5d) and is interpreted to have been derived from rocks of the Alto Guaporé belt, such as granitoids of the Igarapé Enganado and Alto Escondido suites. Four grains yield an age of 1512 Ma, and these grains are derived from rocks of Serra da Providência suite. One grain has an age of 1420 Ma and may have been derived from the tonalitic rocks of the Trinchera complex. The youngest zircon grains (n=2) are most likely derived from granitic leucosome that has an age of 1068 ± 20 Ma. Considering the analytical uncertainty, this age is close to the age of the metamorphic climax of the Nova Brasilândia orogeny (~1110 Ma).

Sample GR-768 (S 13°15'30"/W 60°16'00") is from a slightly elliptical stock that is essentially undeformed. It is a fine-grained equigranular biotite syenogranite, with an elliptical to rounded mafic autolith, an isotropic structure and is intruded into metapsammites of the Aguapeí Group. Its average composition is microcline (36%), oligoclase (27%), quartz (30%), biotite (6%), zircon, apatite and fluorite (<1%). Zircon crystals are prismatic, with aspect ratios 1:2 to 1:4, and are 90-150 µm long and poorly zoned. There are two populations of zircon, zoned magmatic zircon formed at 1025 ± 6 Ma (date of emplacement) (Fig. 5e) and inherited grains with age of the 1569 ± 73 Ma, which is correlated to the Serra da Providência granites. Given the post-tectonic timing implied for the granite, we interpret all of the observed stages of ductile deformation to have occurred prior to 1025 ± 6 Ma.

Sample GR 345 (S 12°05'49"/W 61°25'24") is part of small stock that occur throughout the area and as intrusions in orthogneiss of the São Felipe Suite. This unit was defined by Scandolara and Rizzotto (1998) and occurs near São Felipe city in Rondônia state. The rock is medium- to fine-grained, pink-

colored monzogranite, with an equigranular-to-inequigranular texture. This rock contains biotite and/or hornblende as mafic minerals. Centimeter- to decimeter-wide granular enclaves of granodioritic orthogneiss are commonly observed. This rock is in general massive, but has locally magmatic foliation.

Zircon crystals are subhedral, prismatic grains, clear and of good quality and are 200-250 μm long with broad concentric oscillatory zoning. Of the ten crystals analyzed, nine group tightly, and one is discordant, producing the upper intercept age of 1010 ± 15 Ma (MSDW=0.67) (Fig. 5f). This age is interpreted to be the crystallization age of the granite and is similar to the ages obtained in other late-to post-tectonic granites within the Nova Brasilândia-Aguapeí belt.

DISCUSSION

Main questions regarding the tectonic history and further characterization of the temporal evolution of Nova Brasilândia and Aguapeí belts: (1) Were the belts contiguous, did they form and evolve in the same interval of time, or were they geologically distinct belts? (2) Did these belts form over an already-stabilized crust (craton), or do they represent a suture zone that marks a crustal boundary in the Amazonian craton?

In light of the new geochronological data combined with field data, we suggest a geographical and lithotectonic connection between the Nova Brasilândia and Aguapeí metasedimentary belts. The two belts evolved together and were generated by the reactivation of the Guaporé paleo-suture (Rizzotto and Hartmann 2012, Rizzotto et al. 2013), positioned between the Amazonian craton in the north and the Paraguá Block in the south.

The belts extend up to 900 km inside the cratonic area in SW Amazonian craton (Fig. 1), and their curved shapes depict the ancient Guaporé suture zone.

The Nova Brasilândia belt records the processes of sedimentation, bimodal magmatism and high-grade metamorphism in a rift-basin environment

during a very short time interval from 1122 to 1100 Ma. The orogeny lasted ca. 22 million years from the initial opening of the rift to the final shortening. The youngest zircon group (sample GR-66; 1122 Ma – Fig. 4) of the Nova Brasilândia turbiditic deposits marks the opening of the rift basin, whereas the anatectic melt leucogranitic (sample GR-20; 1110 ± 8 Ma) locally derived from the fusion of the turbidites, indicates the age of peak metamorphism and the closure of basin (Rizzotto 1999). Accordingly, Tohver et al. (2004), using the U-Pb TIMS method, investigated monazite of the Nova Brasilândia metaturbidite, located near sample GR-20, which yielded an age of 1096 ± 5 Ma.

The proto-oceanic-rift phase is marked by the emplacement of gabbros and trondhjemitites prior to 1110 Ma, corresponding to the age of the peak metamorphism affecting these rocks, according to the results obtained from samples GRT-01 to GRT-05.

The Aguapeí belt in turn presents a wider age range, representing the beginning of sedimentation, but shows strong similarities between the ages of the main sources of deposition and the maximum age of sedimentary protholiths when compared with the protolith Nova Brasilândia belt. According to Santos et al. (2005), the deposition of the lower Aguapeí Group (Fortuna Formation) occurred between 1167 ± 27 Ma and 1149 ± 7 Ma. These authors argue that, considering the uncertainties in these ages, the deposition of the sediments of the Fortuna Formation occurred over a period of 46 (maximum) to 12 Ma (minimum). This period is roughly in accordance with the interval associated with the opening and closure of the Nova Brasilândia rift, as noted above. Table II shows the main populations of zircon analyzed to date, in this paper and from previous work.

In the northernmost portion of the Aguapeí belt, the sedimentary rocks are characterized by a gneissic foliation with lower-to-middle amphibolite facies metamorphism and a tight style of folding, and their protoliths are greywackes of the turbiditic

TABLE II
Summary of U-Pb detrital zircon data of the
Nova Brasilândia and Aguapeí Groups.

Sample	Rock	Method	Zircon population (* xenotime)	Younger Zircon (m.a)	References
NM-165	Meta-psammite	U/Pb (shrimp)	1331, 1334, 1342, 1346, 1349, 1354, 1436, 1506, 1513, 1520, 1527,	1331 ± 8	This work
GR-752	phyllite banded	U/Pb TIMS	1390, 1444, 1448, 1449, 1451, 1457, 1465, 1480, 1508, 1650, 1747, 1753, 1764, 1767, 1772, 1788, 1791, 1960 1968, 1991, 2027,	1390 ± 29	This work
NM-182	quartzite	U/Pb TIMS	1334, 1493, 1498, 1499, 1502, 1504, 1545, 1574 1610, 2008	1255 ± 7	This work
AS-88	Meta-siltstone	U/Pb TIMS	1366, 1447, 151, 1617, 1658, 1780, 1894	1299 ± 25	Souza 2011
AS-28	muscovite schist	U/Pb TIMS	1333, 1411, 1521, 1533, 1604, 1607, 1644, 1816	1333 ± 22	Souza 2011
CL-1	Conglomerate	Pb/Pb	1389, 1397, 1432, 1437, 1442, 1496, 1450, 1461, 1463, 1489, 1499, 1507	1319 ± 36	Leite and Saes 2003
MT-1	quartzite	U/Pb (shrimp)	1453, 1420, 1350, 1327, 1271, 1231	1231 ± 14	Santos et al. 2001
JO-02	sandstone	U/Pb (shrimp)	1145*, 1153*	1167 and 1149*	Santos et al. 2005
GR-66	paragneiss	U/Pb (shrimp)	1215, 1320, 1417, 1505, 1599, 1716, 1748, 1904, 2020, 2090	1122 ± 12	This work

environment. Deep-sea turbiditic deposits were also described by Saes (1999) in the Santo Corazón region, where a deep marine turbiditic deposit that comprises deformed immature sedimentary rocks has been deposited into the central rift zone.

These geological features, except for the metamorphic grade, are also found in metaturbidites of the Nova Brasilândia Group (Rizzotto 1999). On the eastern margin of the Serra da Borda, the central portion of the Aguapeí belt, metapelites of the low Fortuna Formation (Souza 2011) have the same structural and lithological characteristics of the northern portion of the Aguapeí belt and have similar detrital zircon ages (Table II). These sedimentary sequences could represent the same lithotectonics unit with little differences in the degrees of deformation and metamorphism and deposited under the same tectonic environments, at the same time. The Aguapeí belt is confined in a narrow aborted rift, whereas the Nova Brasilândia belt was deposited in a proto-oceanic rift environment.

Additionally, providing direct evidence indicating a clear linkage between the two belts, both the Nova Brasilândia belt and Aguapeí belt were involved in a reactivation of the orogenic zone characterized by shear zone of the large-scale sinistral strike-slip motion at ca. 1.0 Ga. This structure was characterized as transcurrent shear zone of the D₂-phase by Rizzotto (1999), and it is recognizable on regional LANDSAT images as large-scale lineaments that extend over hundreds of kilometers. This event is recorded by zircon, titanite and rutile of the mafic rocks (GRT-01, GRT-02 and GRT-05) of the Nova Brasilândia belt (Fig. 4). Tohver et al. (2004) found similar ages in monazite (993 Ma-U/Pb) and hornblende (983 Ma-⁴⁰Ar/³⁹Ar) from the samples collected along the transcurrent shear zone of the Nova Brasilândia belt. Similar ages were found in shear zones that affected the metasedimentary rocks of the Aguapeí Group and neighboring country rocks in the time interval between 1030 and 920 Ma (e.g., Geraldes et al. 1997, Ruiz 2005, Fernandes et al. 2006).

The Sunsás belt, in the southeastern Bolivia shield, presents the same features and geological evolution compared with the Nova Brasilândia-Aguapeí belt. Sedimentary rocks of the Sunsás Group were deposited between 1300 and 1000 Ma. Sunsás deformation show upright NW trending folds formed contemporaneously with low-to-medium grade metamorphism and is structurally well marked by mylonitic shear zones that show preferential sinistral offsets active at ca. 1080-1050 Ma. According to Boger et al. (2005), the stages of ductile deformation must have occurred prior to 1074 Ma, but tectonic reactivation along the shear zones is marked by pegmatite injections (997 ± 3 Ma) and Au-rich zones with molybdenite that yielded a Re/Os age of 994 ± 3 Ma (Isla-Moreno 2009).

On the basis of our findings, we conclude that the Nova Brasilândia, Aguapeí and Sunsás metasedimentary and igneous belts were linked in the late Mesoproterozoic at ca. 1.2 Ga to 0.95 Ga, forming what is here called the Western Amazon belt (WAB). This belt overprints the ca. 1350 crust formed by the Paraguá Block and the basement rocks of the Rondônia-Juruena Province. The main orogenic phase occurred at 1110 Ma (peak metamorphism). The dominant metamorphic grades are high (upper amphibolite to granulite) in the northern part of the belt in Rondônia and low to medium grade (greenschist to lower amphibolite facies) in the eastern zone in Bolivia and in southern Mato Grosso (Sunsás and Aguapeí groups). Reactivation in this belt occurred through extension, uplift and exhumation after ~ 1.0 Ga.

New field geology data coupled with robust geochronology indicate that the Paraguá Block accreted to the southwest margin of Amazonia at ca. 1.47-1.32 Ga (Rizzotto and Hartmann 2012, Rizzotto et al. 2013). Considering this scenario, we suggest that the Nova Brasilândia, Aguapeí and Sunsás groups have a within-plate origin. This scenario is supported by recent paleomagnetic results from 1.42 Ga mafic rocks of the Indiavaí Suite

in Mato Grosso, that demonstrate the latitudinal proximity of the Amazonian craton north of the Nova Brasilândia belt and Guaporé suture zone and the Paraguá craton basement to the south of this suture zone by 1.42 Ga (D'Agrella-Filho et al. 2012). Globally, the tectonic evolution of the Nova Brasilândia, Aguapeí and Sunsás Groups may be related with the final breakup of the supercontinent Columbia, which occurred at about 1.3-1.2 Ga, marked by a major episode of plate-wide extension and rifting, with the emplacement of the gabbros, trondhjemites and volcano-sedimentary rocks. This geological evolution is similar with Laurentia in North America, where occurred the emplacement of the McKenzie, Sudbury, Seal Lake, Harp and Mealy mafic dike swarms and coeval Coppermine River basalts at about 1.3-1.2 Ga (Bartholomew and Hatcher Jr 2010).

Saes et al. (1992) interpreted the terrestrial or shallow marine to deep sedimentary rocks of the Aguapeí and Sunsás groups to be deposited in rift environment. Sedimentary rocks of the Sunsás Group comprise a basal conglomerate overlain successively by two psammitic sequences and an intervening pelitic unit that were deposited in an alluvial-to-deltaic environment (Berrangé and Litherland 1982). Subsequently, Saes et al. (1992) suggested that the deposition of the Sunsás Group has an intra-continental rift environment. These lithotectonic units differ from Nova Brasilândia Group because, in this group, igneous rocks occur temporally and genetically associated with turbidites, reflecting the deposition of these rocks in the rift depocenter. Our theory is that when the reactivation of the Guaporé paleo-suture to generate the Nova Brasilândia-Aguapeí rift, the stretch rate (opening) of the rift, which is now the Nova Brasilândia rift, was high enough to generate a narrow ocean (proto-oceanic rift), while in the other two belts (Aguapeí and Sunsás), there was less opening, less sediment filling, and no ocean development (Fig. 2).

Therefore, over the cratonic area formed by the amalgamation of the Paraguá block and Amazonian craton during the early-middle Mesoproterozoic, the Nova Brasilândia, Aguapeí and Sunsás belts would then reflect zones of intense transpression, within-plate shortening and bimodal mafic-felsic intrusion due to the collisional dynamics, localized along the axis of an intraplate Nova Brasilândia-Aguapeí rift. This tectonic cycle occurred between approximately 1150 and 1100 Ma. Deformation at this time would presumably have been driven by the collision and accretion of the Arequipa-Antofalla basement to the western margin of the Paraguá Block (Loewy et al. 2004). As reported by previously published studies, the Proterozoic tectonic evolution of the Arequipa-Antofalla basement suggests an allochthonous origin, and this microcontinent must have been part of a larger continent before 1.0 Ga. Thus, the Arequipa-Antofalla basement, a Proterozoic crustal block exposed along the modern central Andean margin, is considered to have

accreted to the margin of the Amazonian craton during the Sunsás Orogeny, at approximately 1.2-0.94 Ga (Loewy et al. 2004). According to Dalziel (1994), pervasive metamorphism at 1.2-1.0 Ga in the northern and central domains of the Arequipa-Antofalla basement supports accretion of the Arequipa-Antofalla basement to the Amazonian craton during the Sunsás Orogeny. Therefore, the late Mesoproterozoic orogenic event recorded in the northern and central domains of the Arequipa-Antofalla basement is coeval with events in the southwestern margin of the Amazonian craton. This indicates that the breakup of the supercontinent Columbia was followed rapidly by the assembly of another supercontinent Rodinia, which was assembled along the globally distributed Grenvillian orogens at ca. 1.0 Ga.

We partially agree with the models proposed by Loewy et al. (2004) and Boger et al. (2005) and suggest a tectonic evolution according to the framework depicted in Figure 6.

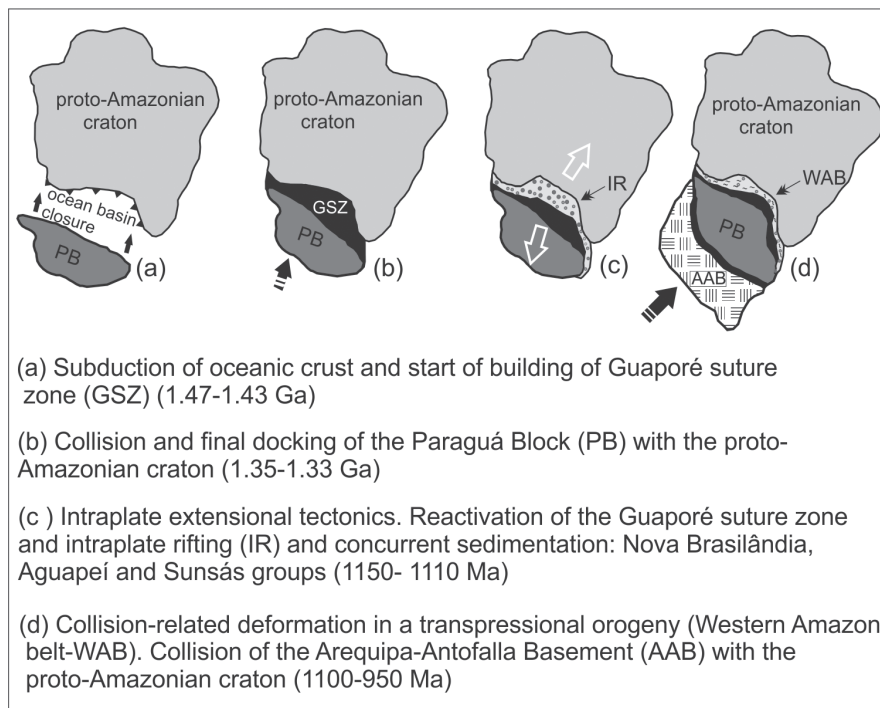


Figure 6 - Hypothetical simplified sketch, showing proposed tectonic evolution of the southwestern Amazonian craton in Mesoproterozoic time (modified of Boger et al. 2005).

CONCLUSIONS

Zircon U-Pb geochronological data and field relationships presented in this contribution are strong evidence for the existence of a continuous intracratonic late Mesoproterozoic orogenic belt (Western Amazon belt) in the southwestern margin of the Amazonian craton.

Considering the final docking of the Paraguá Block and the proto-Amazonian craton during the middle-Mesoproterozoic time (1470-1330 Ma, Rizzotto et al. 2013), the youngest orogenic belt (WAB) of the Amazonian craton is formed by the previously defined Nova Brasilândia, Aguapeí and Sunsás belts, which became interconnected at the end of the Mesoproterozoic, and has a curvilinear shape of approximately 1,000 km in length. The continuity of magnetic anomalies and structural trends suggests that this belt is continuous below the Phanerozoic cover of the Parecis basin.

The opening of the rift that preceded the development of the Nova Brasilândia-Aguapeí belts originated from the reactivation of a paleo-suture (Guaporé suture zone).

We propose that the sedimentary and igneous protoliths for the three orogenic belts exposed in the southwestern margin of the Amazonian craton formed during the late Mesoproterozoic in an intracontinental rift, opening ca. 1150 Ma and closing at 1110 Ma. Therefore, Western Amazon belt's geological history has evolved from the development of a complete Wilson cycle at ca. 40 Ma. Both the Nova Brasilândia-Aguapeí and the Sunsás belts were involved in an orogenic zone characterized by a transtensional stage that was succeeded by a transpressional regime accompanied by crustal shortening that allows the accommodation of N45°E compressional stress field (Rizzotto 1999). This deformation occurred during the accretion of the Arequipa-Antofalla basement to the proto-Amazonian craton.

The latest tectonic movements, before the effective cratonization of the southeastern margin

Amazonian craton, occurred at approximately 1.0 Ga when the reactivation of the previous zone of weakness allowed the development of transcurrent zones, which favored the plutonic activity through the emplacement of the Rio Pardo granites in the Nova Brasilândia region, anorogenic plutons in the Rondônia Tin Province (Bettencout et al. 1999), post-tectonic granite in Sunsás belt (Litherland et al. 1986, Isla-Moreno 2009) and in the Arequipa-Antofalla basement (Wasteneys et al. 1995, Wörner et al. 2000, Loewy et al. 2004). Soon after or during this phase, uplift, exhumation and extension culminated in the generation of intracontinental basins (e.g., Palmeiral Formation and Huanchaca Group).

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RESUMO

Novos dados de U-Pb em zircão e relações de campo, aliados aos dados de levantamento aerogeofísico permitiu-nos sugerir que os Grupos Nova Brasilândia e Aguapeí fazem parte de uma única faixa móvel

constituída por rochas metaígneas e metassedimentares, do final do Mesoproterozóico (1150 Ma -1110 Ma). Esta história geológica é muito semelhante com a evolução intraplaca da Faixa Sunsás, no oriente Boliviano. Assim, propomos que os cinturões Nova Brasilândia, Aguapeí e Sunsás constituem uma única unidade geotectônica (aqui denominada de Faixa Móvel Amazônia Ocidental) desenvolvida no final do Mesoproterozóico, a qual instalou-se num ambiente de rift intracontinental por reativação de uma paleo-sutura (zona de sutura Guaporé). Portanto, a sua história geológica evoluiu a partir do desenvolvimento de um completo Ciclo de Wilson em cerca de 40 Ma. Globalmente, essa evolução tectônica pode estar relacionada com a fragmentação final do supercontinente Columbia. Rochas máficas e trondhjemitos da porção mais setentrional da faixa móvel forneceram idades U-Pb em zircão de ~ 1110 Ma, as quais datam o pico do metamorfismo de alto grau e o fechamento do rift. Isso indica que a fragmentação do supercontinente Columbia foi seguida rapidamente pela montagem do supercontinente Rodínia (~1.1-1.0 Ba) e que a Faixa Móvel Amazônia Ocidental foi construída durante a acreção do Cráton Arequipa-Antofalla ao Cráton Amazônico.

Palavras-chave: Cráton Amazônico, rift intracontinental, Mesoproterozóico, Faixa Móvel Amazônia Ocidental, geocronologia U-Pb.

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SUPPLEMENTARY MATERIAL

Appendix SA – Analytical Procedures