

1 **A novel extraction chromatographic technique for Mo**
2 **separation from geological samples for MC-ICP-MS isotopic**
3 **measurement**

4 Jie Li(1)*, Xi-Rong Liang(1), Li-Feng Zhong(2), Xuan-Ce Wang(3), Zhong-Yuan

5 Ren(1), Sheng-Ling Sun(1), Zhao-Feng Zhang(1), Ji-Feng Xu(1)

6 (1) State Key Laboratory of Isotope Geochemistry, Guangzhou Institute of Geochemistry, Chinese
7 Academy of Sciences, Guangzhou 510640, China

8 (2) Key Laboratory of Marginal Sea Geology, South China Sea Institute of Oceanology, Chinese
9 Academy of Sciences, Guangzhou, 510301, China

10 (3) ARC Centre of Excellence for Core to Crust Fluid Systems (CCFS) and The Institute for
11 Geoscience Research (TIGeR), Curtin University, GPO Box U1987, Perth, 6845, Australia

12

13 Corresponding author. e-mail: jjeli@gig.ac.cn

14

15 We present a novel preconcentration method for the determination of Mo isotope
16 ratios by multicollector-inductively coupled plasma-mass spectrometry (MC-ICP-MS)
17 in geological samples. The method is based on the separation of Mo by extraction
18 chromatography using N-benzoyl-N-phenylhydroxylamine (BPHA) supported on a
19 microporous acrylic ester polymeric resin (Amberlite CG-71). By optimizing the
20 procedure, Mo can be simply and effectively separated from virtually all matrix
21 elements with a single pass through a small volume of BPHA resin (0.5mL). This
22 technique for separation and enrichment of Mo is characterized by high selectivity,
23 column efficiency, and recovery (~100%) and, low total procedural blank (~0.18 ng).
24 A ^{100}Mo - ^{97}Mo double spike is mixed with samples before digestion and column
25 separation, which enables natural mass-dependent isotopic fractionation to be
26 determined with an external precision of $< 0.09\%$ ($\delta^{98/95}\text{Mo}$; 2 sd) by MC-ICP-MS.

27 The mean $\delta^{98/95}\text{Mo}_{\text{SRM3134}}$ (NIST SRM3134 Mo reference standard; Lot No. 891307)
28 composition of the IAPSO seawater reference material measured by this study is 2.00
29 $\pm 0.03\text{‰}$ (2 sd, n = 3), which is consistent with previously published values. The
30 described procedure facilitates efficient and rapid Mo isotopic analysis of various
31 types of geological samples.

32

33 **Keywords:** Molybdenum isotopes; N-benzoyl-N-phenylhydroxylamine; Extraction
34 chromatography; Preconcentration; Double spike; MC–ICP–MS

35

36 The recent advent of multicollector-inductively coupled plasma-mass
37 spectrometry (MC–ICP–MS) has generated increasing interest in isotope fractionation
38 studies of “non-traditional” isotope systems (e.g., Arnold et al. 2004, Bermin et al.
39 2006, Asael et al. 2007, Ripperger and Rehkamper 2007). Of particular interest is the
40 highly redox-sensitive behavior of Mo that is accompanied by significant isotope
41 fractionation, which results in oxic, suboxic, anoxic and euxinic sediments possessing
42 isotopically distinct Mo isotope compositions (e.g., Barling et al. 2001,2004, Siebert
43 et al. 2003, 2006, Anbar 2004). Consequently variations in $\delta^{98/95}\text{Mo}$ can be used as a
44 palaeo-redox proxy to trace the expansion of reducing marine conditions at a number
45 of times throughout Earth's history (e.g., Arnold et al. 2004, Siebert et al. 2005, Pearce
46 et al. 2008). The total range of $\delta^{98/95}\text{Mo}$ compositions observed in terrestrial samples
47 is $\sim 7\text{‰}$. $\delta^{98/95}\text{Mo}$ values can be as low as -3.4‰ (Pearce et al. 2010), as found in
48 hydrothermal waters from Iceland, and as high as $+3.5\text{‰}$ as measured in suboxic

49 sedimentary pore waters (McManus et al. 2002).

50 Given that Mo isotope variations are likely to be on the order of parts per thousand
51 or smaller in magnitude, high-precision analyses are required to quantify the degrees
52 of fractionations and evaluate their origins. However, this presents an analytical
53 challenge because of the difficulties in purifying Mo from potentially interfering ions
54 in geological samples. Moreover, incomplete recovery of Mo could lead to significant
55 isotope fractionation (Anbar et al. 2001, Siebert et al. 2001, Pietruszka et al. 2008). A
56 number of methods have been developed for the separation and purification of Mo
57 from geological samples (e.g., Qi and Masuda 1994, Anbar et al. 2001, Siebert et al.
58 2001, Dauphas et al. 2001, Malinovsky et al. 2005, Pietruszka et al. 2006, 2008,
59 Nakagawa et al. 2008, Pearce et al. 2009). Two-column ion-exchange processes have
60 generally been used for Mo separation from geological samples. This method involves
61 a two-stage separation using an anion-exchange column for separation of Zr and Mn
62 from Mo followed by a cation-exchange column for separation of Fe from Mo.
63 Malinovsky et al. (2005) proposed a single-stage column method using the chelating
64 resin Chelex 100. However, for some geological samples, this method is not
65 sufficiently effective in separating Mo from the Fe and Zr. Thus, a second pass
66 through the same column is necessary in order to remove traces of Zr and to reduce
67 the concentration of Fe to an acceptable level (Malinovsky et al. 2005). Recently,
68 Pearce et al. (2009) developed a single-pass anion-exchange separation method for
69 Mo and Re isotope and abundance determination by MC-ICP-MS from various types
70 of geological samples. Although all these separation and purification methods have

71 been successfully used for Mo isotope determinations, they require relatively large
72 elution volumes and numerous evaporation/transfer steps, which increases sample
73 preparation time and potential blank levels.

74 Yang and Pin (2002) have described a method for preconcentration of Zr, Hf, Nb
75 and Ta in rock samples using N-benzoyl-N-phenylhydroxylamine (BPHA) supported
76 on a microporous acrylic ester polymeric resin (Amberlite CG-71). BPHA is relatively
77 small in size, has abundant π electrons, is stable in highly acidic solutions, and acts as
78 a superior ligand for solvent extraction of group 4, 5 and 6 elements in the Periodic
79 Table, such as Nb, Ta, Zr, Hf, Ti, Mo and W (Caletka and Krivan 1989, Goguel 1992,
80 Yang and Pin 2002, Shinotsuka and Suzuki 2007, Li et al. 2010). Amberlite CG-71
81 resin possesses excellent chemical and physical stability, and can absorb large
82 amounts of BPHA because it contains relatively large numbers of active aromatic sites
83 that allow π - π interactions. BPHA extraction chromatographic resin has high
84 selectivity for the adsorption of tetra-, penta- and hexavalent elements, whereas other
85 matrix elements have little affinity for the resin (Yang and Pin 2002).

86 This study reports the first utilization of BPHA extraction chromatographic resin
87 for the pre-concentration of Mo in geological samples and for the determination of
88 Mo isotopic ratios by MC-ICP-MS. The chemical separation parameters were
89 systematically investigated and optimized. It was found that Mo has very strong
90 affinity to this resin even in 0.1 mol L⁻¹ HF, and 6 mol L⁻¹ HF was necessary to
91 achieve quantitative recovery of Mo. The key objective of this study was to develop a
92 chemical separation procedure for subsequent accurate and precise determination of

93 $\delta^{98/95}\text{Mo}$ and Mo concentrations by ^{100}Mo – ^{97}Mo double-spike MC-ICP-MS analysis.

94 The efficiency and widespread applicability of our technique is demonstrated by
95 replicate analyses of geological reference materials and the IAPSO seawater standard.

96 **Experimental methods**

97 **Reagents and materials**

98 The Johnson Matthey Specpure Mo plasma standard solution (Lot No. 012773) and
99 NIST SRM3134 (Lot No. 891307) were used as the Mo isotopic standards. A
100 ^{97}Mo – ^{100}Mo double spike was prepared as a mixture of individual concentrated spike
101 solutions of isotopically enriched ^{97}Mo (Batch 159791) and ^{100}Mo (Batch 159992)
102 which were purchased from the Oak Ridge National Laboratory, USA. The
103 compositions of the ^{97}Mo – ^{100}Mo double spike and Mo reference solution were
104 calibrated by doping with Pd using the method detailed by Siebert et al. (2001) (for
105 double-spike and standard calibration results see Table 1). BPHA (Analytical grade,
106 Aladdin Reagent Inc. China) and Amberchrom CG-71 (50–100 μm particle size,
107 Supelco Inc., USA) were used without further purification. A synthetic multi-element
108 solution with a concentration of 1 $\mu\text{g mL}^{-1}$ was prepared from AccuStandard fluoride
109 soluble elements (MISA 05-1), rare earth metals (MISA 01-1), transition metals
110 (MISA 06-1), alkaline earth elements (MISA 04-1), and precious metals (MISA 02-1)
111 for use in optimizing the column separation parameters. HCl, HF and HNO_3 were
112 purified by DST-1000 sub-boiling stills (Savillex Corporation, USA). Ultrapure water
113 prepared with a Milli-Q system (Millipore) was used to dilute all the acids, and is
114 referred to as Milli-Q water hereafter in this paper.

115 **Preparation of BPHA resin**

116 We prepared the BPHA resin column using procedures modified from those
117 described by Yang and Pin (2002). Five grams of Amberchrom CG-71 resin was
118 impregnated with 40 mL of 3% BPHA solution in ethanol rather than chloroform. The
119 mixture was shaken at room temperature (25 °C) for 2 h to ensure complete saturation
120 and was then stored in a refrigerator without removing the ethanol prior to use. This
121 material is hereafter referred to as the BPHA resin. Approximately 0.5 mL of BPHA
122 resin (ca. 0.2 g of dry weight) was packed in a Bio-Spin chromatography column
123 (Bio-Rad Laboratories Inc., USA).

124 **Sample digestion and preparation**

125 Sample preparation and all chemical separations and evaporations were carried out
126 in a clean environment (Class 1000). Approximately 40–70 mg of samples powder
127 was weighted accurately and combined with the ^{97}Mo – ^{100}Mo double spike in 15 mL
128 PFA beakers. The beakers were then placed on a hotplate at 120 °C along with ~ 6 mL
129 of a 2:1 mixture of HF (22 mol L⁻¹) and HNO₃ (14 mol L⁻¹). To achieve total sample
130 dissolution, the PFA beakers were placed in an ultrasonic bath for 5 min at various
131 intervals during the dissolution period. After digestion and dryness at 120 °C, the
132 samples were dissolved in 1 mL of concentrated HCl, and again evaporated to dryness.
133 The residue was re-dissolved in 2–4 mL of a mixture of 0.1 mol L⁻¹ HF/1mol L⁻¹ HCl,
134 at which point it was ready for column separation. Approximately 10 mL of the
135 seawater sample was acidified with 1 mL of concentrated HCl, and the ^{97}Mo – ^{100}Mo
136 double spike was then added. The spiked and acidified seawater solution was then

137 evaporated to dryness before being re-dissolved in 4 mL of 0.1 mol L⁻¹ HF/1mol L⁻¹
138 HCl.

139 **Chemical separation**

140 Prior to use, the BPHA resin column was washed with 6 mL of 6 mol L⁻¹ HF/1 mol
141 L⁻¹ HCl and 4 mL of Milli-Q water. Prior to sample loading, the column was
142 conditioned with 2 mL of 0.1 mol L⁻¹ HF/1mol L⁻¹ HCl. The chemical separation steps
143 [and along with previously published methods](#) were outlined in Table 2. A 2 mL aliquot
144 of the sample solution was loaded onto the column and the resin was washed with 8
145 mL of 0.1 mol L⁻¹ HF/1mol L⁻¹ HCl. This step effectively removes matrix and
146 interfering elements (Fe, Mn, Zr, Ru, Cu, Zn, Ni, etc.) from the sample. Finally, the
147 adsorbed Mo was eluted by passing in 8 mL of 6 mol L⁻¹ HF/1mol L⁻¹ HCl. The Mo
148 was collected in 15 mL PFA vials and evaporated on a hot plate at 120 °C to dryness.
149 Three drops of concentrated HNO₃ and H₂O₂ were added to the evaporated Mo to
150 decompose any organic residue. Following this, 1 mL of 3% HNO₃ was added to the
151 Mo residue, after which the solution was ready for Mo isotopic ratio measurement by
152 MC-ICP-MS.

153 **Mass spectrometry measurements**

154 During the method development stage of this work, a quadrupole ICP-MS
155 (Thermo-Scientific Xseries-2) was used in the conventional mode for
156 semi-quantitative and quantitative elemental measurements. Typical operating
157 conditions for this instrument have been reported elsewhere (Li et al., in press).

158 Mo isotopic ratios were determined on a Thermo-Fisher Scientific Neptune Plus

159 MC–ICP–MS at the State Key Laboratory of Isotope Geochemistry, Guangzhou
160 Institute of Geochemistry (GIG), Chinese Academy of Sciences (CAS), Guangzhou,
161 China. This instrument is equipped with eight moveable Faraday collectors and one
162 fixed central collectors, which are linked to amplifiers with 10^{11} Ω resistors. Samples
163 and standards were introduced into the plasma through an Aridus II® desolvating
164 sample introduction system (CETAC Technologies, Omaha, USA). This system
165 removes water from the sample solutions by passage through a desolvating membrane
166 at 160 °C, providing a “dry” sample to the plasma, which increases sensitivity by a
167 factor of 5~10 over wet plasma sample-introduction systems. The typical Mo
168 sensitivity achieved was 180–200 V ppm⁻¹. The operating conditions and Faraday
169 collector configuration for the Neptune Plus are given in Table 3. Data for all samples
170 and standards were acquired in 6 blocks of 10 ratios. Subsequently, these 60 ratios
171 were filtered with a one-pass 2sd outlier removal test, and the mean value and
172 standard error (SE) were calculated. After each analysis, the sample was sequentially
173 washed from the system for ten min using solutions of 3% HNO₃, 0.2 mol L⁻¹ HF/3%
174 HNO₃ and 3% HNO₃. The residual Mo signal was typically < 1 mv, which is
175 insignificant to the ion beams measured during sample analysis.

176 Prior to starting an analytical session, the reference solution (NIST SRM3134) was
177 typically analyzed five times to check the MC–ICP–MS performance. During an
178 analytical session, the reference solution was measured after each batch of three
179 samples. Correction for mass fractionation during chemical separation, as well as
180 during mass spectrometry followed the double–spike deconvolution methods

181 described by Siebert et al. (2001).

182 Mo isotopic ratios of all samples were normalized to those of the NIST SRM3134

183 Mo standard and expressed in conventional δ notation:

$$184 \quad \delta^{98/95}\text{Mo}(\text{‰}) = [({}^{98}\text{Mo}/{}^{95}\text{Mo})_{\text{sample}} / ({}^{98}\text{Mo}/{}^{95}\text{Mo})_{\text{SRM3134}} - 1] \times 1000$$

185 **Results and discussion**

186 **Adsorption capacity of BPHA resin**

187 The adsorption capacity or maximum amount of Mo adsorbed per gram of freshly
188 synthesized BPHA resin was determined. The adsorption capacity of the BPHA resin
189 is defined as the amount of metal ions that can be extracted per unit mass under the
190 optimized chemical separation conditions. The method used here is based on that
191 recommended by Maquieira et al, (1994). Solutions containing $10 \mu\text{g g}^{-1}$ of Mo were
192 passed through the column continuously until detectable signal of Mo was observed in
193 the eluent. The maximum volume (V) of solution at concentration (c) and the amount
194 of metal adsorbed on the column were then calculated by the following equation:

$$195 \quad C = cV/W$$

196 where C is the adsorption capacity and W is the weight of the BPHA resin. In this
197 study, the adsorption capacity for Mo was found to be 0.72 mg g^{-1} . The resin exchange
198 capacity was reduced by ~40 % after the resin had been used once. However, the
199 BPHA column could be refreshed by eluting a 3 % BPHA solution in ethanol through
200 it.

201 **Optimization of the chemical separation procedure**

202 The presence of fluorides has a strong negative effect on the extraction of Zr, Hf,

203 Nb, Ta, Mo and W by BPHA (Caletka and Krivan, 1989, Goguel, 1992). We
204 examined the effect of HF concentration (0.1 to 8 mol L⁻¹ HF) on the collection of Mo
205 by BPHA resin column and found that Ti, Zr, Hf, Nb and W were hardly retained on
206 BPHA resin at low concentrations of HF (0.1 mol L⁻¹ HF), whereas the behaviors of
207 Ta and Mo on the resin were unaffected. This observation is consistent with those of
208 Goguel (1992), in that Mo and Ta have a greater affinity for BPHA. Mo and Ta are
209 gradually eluted from the column if the concentration of HF is above 1 mol L⁻¹, and
210 **are quantitatively striped when the HF concentration is > 5 mol L⁻¹**. The effect of HCl
211 concentration on the collection of Mo from the BPHA resin column was also
212 examined. In all the cases, from 0.5 to 4 mol L⁻¹ HCl, Ti, Zr, Hf, Nb, Ta, W and Mo
213 are completely absorbed on the BPHA resin column, whereas other elements pass
214 straight through the column. In subsequent experiments, Mo collection was carried
215 out with 0.1 mol L⁻¹ HF/1 mol L⁻¹ HCl as matrix elements such as Ti, Nb, Zr, Hf and
216 W are not retained on the BPHA resin in this acid mix and Mo can be adsorbed by a
217 small resin volume. The Mo adsorbed on the column was removed in 8 mL of 6 mol
218 L⁻¹ HF/1 mol L⁻¹HCl. Figure 1 shows a typical elution profile for Fe, Mn, Cr, Ti, Rb,
219 Sr, La, Zr, Hf, Nb, Ta, Ru and Mo. It is apparent that no Fe, Mn, Cr, Ru and rare earth
220 elements, as well as Ti, Zr, Hf, Nb and W, are extracted by the BPHA resin in a 0.1
221 mol L⁻¹ HF/1 mol L⁻¹ HCl acid mixture. Only Ta is quantitatively extracted with Mo.

222 **Isotopic fractionation during BPHA resin chemical separation**

223 It is widely recognized that Mo isotope fractionation can result from incomplete
224 Mo recovery during ion chromatographic separation (e.g., Anbar et al. 2001, Siebert

225 et al. 2001, Pietruszka and Reznik 2008). Isotopic fractionation during chemical
226 separation of Mo was examined using the BPHA resin column as described by Pearce
227 et al. (2009). Solution aliquots of 2 mL containing 0.5429 ppm Mo in 0.1 mol L⁻¹
228 HF/1 mol L⁻¹ HCl were prepared from the NIST SRM3134 Mo standard solution and
229 passed through the BPHA column. Three of these solutions had an appropriate
230 amount of the ¹⁰⁰Mo–⁹⁷Mo double spike (the ⁹⁷Mo_{spike}/⁹⁷Mo_{sample} = ~ 3) added prior to
231 separation (“spiked”) whereas the other three solutions had the spike added afterwards
232 (“unspiked”). The “spiked” and the “unspiked” solutions gave mean δ^{98/95}Mo values
233 of 0.01 ± 0.05‰ (2 sd, n = 3) and 0.03 ± 0.02‰ (2 sd, n = 3), respectively (Table 4).
234 These determinations are indistinguishable within analytical uncertainty, and are also
235 the same as the expected δ^{98/95}Mo composition (0.00 ± 0.09‰; Fig. 2). Furthermore,
236 quantitative recovery of Mo from the column was confirmed by assessing the Mo
237 yields from each solution (Table 4). Measured Mo recoveries for the “unspiked” and
238 “spiked” solution are 98.9 ± 1.1% (2 sd; n = 3) and 99.4 ± 0.7% (2 sd; n = 3),
239 respectively. At this good recovery rate, an isotope fractionation of Mo by BPHA
240 column separation is unlikely.

241 **Matrix effect**

242 Coexisting elements in solutions can interfere with Mo isotopic ratio
243 measurements and produced inaccurate results. Although the BPHA resin column
244 pre-concentration method effectively separated the matrix and interfering elements
245 from Mo, Ta is collected along with Mo (Fig. 1). To test the matrix effect from Ta, we
246 measured Mo isotopic ratios on NIST SRM3134 Mo standard solutions doped with

247 various amounts of Ta. (Fig. 2). No significant change in measured $\delta^{98/95}\text{Mo}$ was
248 found even at ratios of Ta/Mo = 100. Ta/Mo ratios in most geological samples are \leq
249 100 and, in particular, sedimentary rocks are highly depleted in Ta. Thus, we
250 concluded that potential matrix effects from the presence of Ta are insignificant.
251 Isobaric interferences from elemental and molecular ions in the solution may occur
252 during Mo isotopic ratio measurements. Zr interferes with ^{92}Mo , ^{94}Mo , and ^{96}Mo , and
253 Ru interferes with ^{96}Mo , ^{98}Mo , and ^{100}Mo . Polyatomic ions such as $^{52}\text{Cr}^{40}\text{Ar}^+$,
254 $^{55}\text{Mn}^{40}\text{Ar}^+$, $^{56}\text{Fe}^{40}\text{Ar}^+$, $^{60}\text{Ni}^{40}\text{Ar}^+$, and $^{80}\text{Se}^{16}\text{O}^+$ can also possibly interfere on Mo
255 isotopes. However, these interferences are completely removed from the separated
256 Mo by BPHA resin column, and concentrations of Zr, Cr, Mn, Fe, Ni, and Se in the
257 Mo solution are very low (a few ppb), and Ru was not detected in the analyte solution
258 by ICP-MS in the separated Mo solution. Moreover, the Aridus II® desolvating
259 sample-introduction system reduces these polyatomic ions to negligible level and,
260 coupled with the efficiency of our Mo separation procedures, means that our Mo
261 isotope data are unaffected by inference.

262 **Precision and accuracy**

263 The precision of our Mo isotope determination was assessed by [analyzing spiked](#)
264 [aliquots of two](#) Mo standard solutions: NIST SRM3134 and the Johnson Matthey
265 Specpure Mo plasma standard solution (JMC). It should be noted that the lack of a
266 commonly used Mo isotopic standard, means that the standards used by different
267 laboratories appear to be isotopically different at the level of $\sim 0.2\%$. [Recently, Wen et](#)
268 [al. \(2010\) and Goldberg et al. \(2013\) recommend that NIST SRM3134 Mo standard](#)

269 should be used as the universal reference material for reporting Mo isotopic
270 composition of natural samples. Here, we adopt NIST SRM3134 as the reference
271 standard and compare our $\delta^{98/95}\text{Mo}_{\text{SRM3134}}$ values with previously reported values (Fig.
272 3). The average $\delta^{98/95}\text{Mo}_{\text{SRM3134}}$ for NIST SRM3134 and JMC were $0.00 \pm 0.09\%$ (2
273 sd, $n = 21$) and $-0.20 \pm 0.06\%$ (2 sd, $n = 14$), respectively. The accuracy of our Mo
274 isotope determination was also assessed by analyzing the IAPSO Atlantic seawater.
275 The measured $\delta^{98/95}\text{Mo}_{\text{SRM3134}}$ value of IAPSO was $2.00 \pm 0.03\%$ (Table 5), which is
276 comparable to the value reported by Greber et al. (2012).

277

278 **Application to geological materials**

279 The applicability of our technique for analysis of natural samples was assessed by
280 replicate analysis of three different types of reference samples: offshore marine
281 sediment GBW07316, the USGS basalt BHVO-2, and IAPSO Atlantic seawater
282 (Table 5).

283 The average $\delta^{98/95}\text{Mo}_{\text{SRM3134}}$ value determined for the BHVO-2 ($n = 3$) was $-0.05 \pm$
284 0.11% (2 sd) with individual analyses varying from -0.11% to $+0.01\%$. This mean
285 value for BHVO-2 is similar to values reported for this standard by Pearce et al.
286 (2009), and also falls within the range of other igneous rocks (Siebert et al. 2003). The
287 average Mo concentration determined for this basalt was $3.9 \pm 2.3 \mu\text{g g}^{-1}$ (2 sd) with
288 individual analyses varying from 2.84 to $5.15 \mu\text{g g}^{-1}$. Our average Mo abundance for
289 BHVO-2 is identical to the value reported by Hu et al. (2008). We note that the
290 measured $\delta^{98/95}\text{Mo}_{\text{SRM3134}}$ values for BHVO-2 were positively correlated with Mo

291 abundances, which suggests that the inhomogeneous distribution of minor magmatic
292 sulfides is responsible for these covariations (Voegelin et al. 2012). The mean
293 $\delta^{98/95}\text{Mo}_{\text{SRM3134}}$ composition of the analysed GBW07316 offshore marine sediment
294 reference material is $-0.60 \pm 0.10\text{‰}$, which falls within the range of Mo isotope
295 values obtained for pelagic clays and Fe–Mn crust (Siebert et al. 2003). However, it
296 should be noted that the $\delta^{98/95}\text{Mo}$ composition of such samples can vary considerably
297 depending on paleo-redox conditions (Siebert et al. 2003). The measured Mo
298 concentration for the GBW07316 is $5.32 \pm 0.42 \mu\text{g g}^{-1}$ (2 sd), which is consistent with
299 the reference value for this standard when analytical uncertainties are considered. The
300 measured $\delta^{98/95}\text{Mo}_{\text{SRM3134}}$ value of seawater obtained by analysis of IAPSO was 2.00
301 $\pm 0.03\text{‰}$ which is similar to the value reported by Greber et al. (2012). The
302 concentration of Mo in IAPSO was $0.010 \mu\text{g g}^{-1}$ which is also indistinguishable from
303 previously studies (e.g., Siebert et al. 2003, Pearce et al. 2009, Nakagawa et al. 2008,
304 Grebert et al. 2012).

305 In summary, these results confirm that our analytical procedure is suitable for
306 precise and accurate determination of Mo isotopic composition and Mo abundances in
307 a wide range of geological materials.

308 **Total procedural blanks**

309 The procedural blank was determined using exactly the same methods applied to
310 the natural samples. Total procedural blanks varied from 0.14 ng to 0.21 ng (mean =
311 0.18 ± 0.06 ng; 2 sd; n = 3). Our total procedural blanks are considerably lower than
312 those of reported for other Mo chemical separation procedures (i.e., 0.5 to 47 ng; e.g.,

313 Siebert et al. 2001, Pearce et al. 2009, Nakagawa et al. 2012, Voegelin et al. 2012,
314 Goldberg et al. 2013), due to the relatively small volume of resin and acids used by
315 our method. The total Mo blank contribution never amounts > 0.3% of the sample Mo
316 and, as such, has no effect on the measured Mo isotope composition of the samples.

317 **Conclusions**

318 We have developed a new method using BPHA extraction chromatography and
319 MC–ICP–MS for the determination of Mo isotopic compositions in geological
320 materials. The method has excellent selectivity and efficiency. This single-pass
321 technique is relatively straightforward and has ~100% Mo recovery, which ensures
322 that Mo isotope fractionation does not occur during chemical separation of Mo. The
323 small volume of the resin (0.5 mL) and eluates (<20 mL) required results in very low
324 total procedural blanks (~0.18 ng). The BPHA method thus offers significant
325 analytical advantages over some of the other chemical procedures that are currently
326 used for Mo separation and isotope determination. However, a disadvantage of our
327 method is the relatively large volumes of HF used, and the highly toxic and corrosive
328 nature of even dilute HF acid. Multiple analyses of marine sediment reference
329 material GBW07316, basalt reference material BHVO-2, and IAPSO seawater salinity
330 reference materials demonstrate that this new analytical procedure can be reliably
331 applied to a wide range of geological materials.

332 **Acknowledgements**

333 Dr. Yuexing Feng (The University of Queensland) and Dr. Aaron (Stallard Scientific
334 Editing, New Zealand) are acknowledged for English editing. We are grateful to Tian

335 You Rong (Thermo Fisher Scientific Inc. China) and Zhang Le (State Key Laboratory
336 of Isotope Geochemistry, Guangzhou Institute of Geochemistry, China) for their
337 assistance with MC–ICP–MS measurements. Dr. Ma Jin-Long (State Key Laboratory
338 of Isotope Geochemistry, Guangzhou Institute of Geochemistry, China) kindly
339 provided the IAPSO seawater. Editor Prof. Bill McDonough and [three anonymous](#)
340 [reviewers are thanked for their thoughtful and constructive comments](#) /editorial advice
341 that significantly improved this manuscript. This study was jointly supported by the
342 Technology Development Project (SKLIG-JY-12-03) from the State Key Laboratory
343 of Isotope Geochemistry and the National Science Foundation of China (Grants
344 41173038 and 41203010), as well as an ARC ESTART Fellowship through the ARC
345 Centre of Excellence for Core to Crust Fluid Systems (CCFS). This is contribution No.
346 IS-1777 from GIGCAS, TIGeR publication No. 496 and CCFS contribution 369.

347

348

349

350

351

352

353

354

355

356

- 357 **References**
- 358 **Anbar A.D., Knab K.A. and Barling J. (2001)**
- 359 Precise determination of mass-dependent variations in the isotopic composition of molybdenum
- 360 using MC-ICP-MS. *Analytical Chemistry*, 73, 1425-1431.
- 361
- 362 **Anbar A.D. (2004)**
- 363 Molybdenum stable isotopes: Observations, interpretations and directions. In: Johnson C.M.,
- 364 Beard B. and Albarède F. (eds), *Geochemistry of non-traditional stable isotopes. Reviews in*
- 365 *Mineralogy and Geochemistry*, 55, 429-454.
- 366
- 367 **Arnold G.L., Anbar A.D., Barling J. and Lyons T.W. (2004)**
- 368 Molybdenum isotope evidence for widespread anoxia in mid-Proterozoic oceans. *Science*, 304,
- 369 87-90.
- 370
- 371 **Arnold G. L., Weyer S. and Anbar A. D. (2004)**
- 372 Iron isotope variations in natural materials measured using high mass resolution MC-ICP-MS.
- 373 *Analytical Chemistry*, 76, 322-327.
- 374
- 375 **Asael, D., Matthews, A., Bar-Matthews, M. and Halicz, L. (2007)**
- 376 Copper isotope fractionation in sedimentary copper mineralization (Timna Valley,
- 377 Israel). *Chemical Geology*, 243, 238-254
- 378
- 379 **Barling J., Arnold G. L. and Anbar A. D. (2001)**
- 380 Natural mass-dependent variation in the isotopic composition of molybdenum. *Earth Planet. Sci.*
- 381 *Lett.* 193, 447-457.
- 382
- 383 **Barling J. and Anbar A. D. (2004)**
- 384 Molybdenum isotope fractionation during adsorption by manganese oxides. *Earth and Planetary*
- 385 *Science Letters*, 217, 315-329.
- 386
- 387 **Bermin J., Vance D., Archer C. and Statham P.J. (2006)**
- 388 The determination of the isotopic composition of Cu and Zn in seawater. *Chemical*
- 389 *Geology*, 226, 280-297.
- 390
- 391 **Caletka R. and Krivan V. (1989)**
- 392 Extraction of molybdenum and tungsten with various reagents. *Fresenius' Zeitschrift für*
- 393 *analytische Chemie*, 332, 866-873.
- 394
- 395 **Dauphas N., Reisberg L. and Marty B. (2001)**
- 396 Solvent extraction, ion chromatography, and mass spectrometry of molybdenum isotopes.
- 397 *Analytical Chemistry*, 73, 2613-2616.
- 398
- 399 **Goguel R. (1992)**

400 Group separation by solvent extraction from silicate rock matrix of Nb, Mo, Ta, and W at low
401 levels for ICP-MS. *Fresenius' Journal of Analytical Chemistry*, 344, 326-333.

402 **Goldberg T., Gordon G., Izon G., Archer C., Pearce C. R., McManus J., Anbar A. D. and**
403 **Rehkämper M. (2013).**
404 Resolution of inter-laboratory discrepancies in Mo isotope data: an intercalibration. *Journal of*
405 *Analytical Atomic Spectrometry*, 28, 724–735.

406

407 **Greber N. D., Siebert C., Nägler T. F. and T. Pettke. (2012)**
408 $\delta^{98/95}\text{Mo}$ values and Molybdenum Concentration Data for NIST SRM 610, 612 and 3134:
409 Towards a Common Protocol for Reporting Mo Data. *Geostandard and Geoanalytical Reserche*,
410 36, 291-300.

411

412 **Hu Z.C. and Gao S. (2008)**
413 Upper crustal abundances of trace elements: A revision and update. *Chemical Geology*, 253,
414 205-221.

415 **Li J., Zhong L.F., Tu X.L., Liang X.R. and Xu J.F. (2010)**
416 Determination of rhenium content in molybdenite by ICP-MS after separation of the major matrix
417 by solvent extraction with N-benzoyl- N-phenylhydroxalamine. *Talanta*, 81, 954–958.

418

419 **Li J., Jiang X.Y., Xu J.F., Zhong L.F., Wang X.C., Wang G.Q. and Zhao P.P. (2013)**
420 Determination of platinum-group elements and Re-Os isotopes using ID-ICP-MS and N-TIMS
421 from a single digestion after two-stage column separation. *Geostandard and Geoanalytical*
422 *Research*, (doi: 10.1111/j.1751-908X.2013.00242.x).

423

424 **Malinovsky D., Rodushkin I., Baxter D.C., Ingri J. and Ohlander B. (2005)**
425 Molybdenum isotope ratio measurements on geological samples by MC-ICP-MS. *International*
426 *Journal of Mass Spectrometry*, 245, 94-107.

427

428 **Maquiera A., Elmahadi H. and Puchades R. (1994)**
429 Immobilized cyanobacteria for on-line trace metal enrichment by flow injection atomic absorption
430 spectrometry. *Analytical Chemistry*, 66, 3632-3638.

431

432 **McManus J., Nägler T.F., Siebert C., Wheat C.G. and Hammond D.E. (2002)**
433 Oceanic molybdenum isotope fractionation: Diagenesis and hydrothermal ridge-flank alteration.
434 *Geochemistry, Geophysics, Geosystems*, 3, 1078, doi:10.1029/2002 GC00035.

435

436 **Nakagawa Y., Firdaus M.L., Norisuye K., Sohrin Y., Irisawa K. and Hirata T. (2008)**
437 Precise isotopic analysis of Mo in seawater using multiple collector-inductively coupled mass
438 spectrometry coupled with a chelating resin column preconcentration method. *Analytical*
439 *Chemistry*, 80, 9213-9219.

440

441 **Pearce C.R., Cohen A.S., Coe A.L. and Burton K.W. (2008)**
442 Molybdenum isotope evidence for global ocean anoxia coupled with perturbations to the carbon

443 cycle during the Early Jurassic. *Geology*, 36, 231-234.
444

445 **Pearce C.R., Cohen A.S. and Parkinson, I.J. (2009)**
446 Quantitative separation of molybdenum and rhenium from geological materials for isotopic
447 analysis by MC-ICP-MS. *Geostandard and Geoanalytical Research*, 33, 219-229.
448

449 **Pearce C.R., Burton K.W., Pogge von Strandmann P.A.E., James., R.H., and Gislason S.R.**
450 **(2010)**
451 Molybdenum isotope behaviour accompanying weathering and riverine transport in a basaltic
452 terrain, *Earth and Planetary Science Letters*, 295, 104-114
453

454 **Pietruszka A.J., Walker R.J. and Candela P.A. (2006)**
455 Determination of mass-dependent molybdenum isotopic variations by MC-ICP-MS: An evaluation
456 of matrix effects. *Chemical Geology*, 225, 121-136.
457

458 **Pietruszka A.J. and Reznik A.D. (2008)**
459 Identification of a matrix effect in the MC-ICP-MS due to sample purification using ion exchange
460 resin: An isotopic case study of molybdenum. *International Journal of Mass Spectrometry*, 270,
461 23-30.
462

463 **Qi L. and Masuda A. (1994)**
464 The isotopic composition and atomic-weight of molybdenum. *International Journal of Mass*
465 *Spectrometry and Ion Processes*, 130, 65-72.
466

467 **Ripperger S and Rehkaemper M. (2007)**
468 Precise determination of cadmium isotope fractionation in seawater by double spike
469 MC-ICPMS, *Geochimica et Cosmochimica Acta*, 71, 631-642.
470

471 **Shinotsuka K and Suzuki K. (2007)**
472 Simultaneous determination of platinum group elements and rhenium in rock samples using
473 isotope dilution inductively coupled plasma mass spectrometry after cation exchange separation
474 followed by solvent extraction. *Analytica Chimica Acta*, 603, 129-139.
475

476 **Siebert C., Nagler T.F., von Blanckenburg F. And Kramers J.D. (2003)**
477 Molybdenum isotope records as a potential new proxy for paleoceanography. *Earth and Planetary*
478 *Science Letters*, 211, 159-171.
479

480 **Siebert C., McManus J., Bice A., Poulson R. and Berelson W.M. (2006)**
481 Molybdenum isotope signatures in continental margin marine sediments. *Earth and Planetary*
482 *Science Letters*, 241, 723-733.
483

484 **Siebert C., Kramers J.D., Meisel T., Morel P. and Nagler T.F. (2005)**
485 PGE, Re-Os, and Mo isotope systematics in Archean and early Proterozoic sedimentary systems
486 as proxies for redox conditions of the early Earth. *Geochimica et Cosmochimica Acta*, 69,

487 1787-1801.

488

489 **Siebert C., Nägler T.F. and Kramers J.D. (2001)**

490 Determination of molybdenum isotope fractionation by double-spike multicollector inductively
491 coupled plasmamass spectrometry. *Geochemistry Geophysics Geosystems*, 2, 2000GC000124.

492

493 **Voegelin A. R., Nagler T.F., Pettke T., Neubert N., Steinmann M., Pourret O. and Villa I. M.**
494 **(2012)**

495 The impact of igneous bedrock weathering on the Mo isotopic composition of stream waters:
496 Natural samples and laboratory experiments. *Geochimica et Cosmochimica Acta*, 86, 150-165.

497

498 **Wen H., Carignan J., Cloquet C., Zhu X. and Zhang Y. (2010)**

499 Isotopic delta values of molybdenum standard reference and prepared solutions measured by
500 MC-ICP-MS: Proposition for delta zero and secondary references. *Journal of Analytical Atomic*
501 *Spectrometry*, 25, 716–721.

502

503 **Yang X.J. and Pin C. (2002)**

504 Determination of niobium, tantalum, zirconium and hafnium in geological materials by extraction
505 chromatography and inductively coupled plasma mass spectrometry. *Analytica Chimica Acta*, 458,
506 375–386.

507

508

509

510

511

512

513

514

515

516

517

518

519

520

521

522

523

524

525

526

527

528

529

530

531 **Table captions**

532 Table 1. Compositions of the ^{100}Mo – ^{97}Mo double spike and NIST SRM3134 Mo
 533 standard.

	$^{100}\text{Mo}/^{97}\text{Mo}$	$^{98}\text{Mo}/^{97}\text{Mo}$	$^{96}\text{Mo}/^{97}\text{Mo}$	$^{95}\text{Mo}/^{97}\text{Mo}$	$^{94}\text{Mo}/^{97}\text{Mo}$	$^{92}\text{Mo}/^{97}\text{Mo}$
Double spike	0.979966	0.061740	0.027146	0.018081	0.009581	0.016851
2s (n=10)	0.000075	0.000074	0.000057	0.000056	0.000031	0.000055
Standard solution	1.003817	2.523671	1.747802	1.671446	0.972042	1.564659
2s (n=10)	0.000060	0.000171	0.000132	0.000127	0.000095	0.000113

534

535 Table 2. The BPHA extraction chromatographic column procedure for Mo purification
 536 chemistry, and along with other purification methods for comparison.

Mehods	Resin	Volume (mL)	Step	Volume(mL)	Purpose
This study	BPHA resin	0.5	0.1 M HF/1 M HCl	2.0	Condition
			0.1 M HF/1 M HCl	2–4	Load
			0.1 M HF/1 M HCl	8	Matrix removal
			6 M HF/1 M HCl	8	Collection of Mo
Two column ion-exchange (Pietruszka et al. 2006)	AG1X8	10	6M HCl	30	Condition
			6M HCl	5	Loading
			6M HCl	40	Matrix removal
			0.1 M HF/0.01 M HCl	40	Matrix removal
			1 M HCl	120	Collection of Mo
			5 M HNO ₃	50	Collection of Mo
			1.4 M HCl	20	Condition
AG 50WX8	3.5	1.4 M HCl	10	Load and Collection of Mo	
Single column chelating resin (Malinovsky et al. 2005)	Chelex 100	3	0.3 M HNO ₃	10	Condition
			0.3 M HNO ₃	12	Load
			0.07 M HCl	10	Matrix removal
			0.1 M HF	50	Matrix removal
			6 M NH ₃	12	Collection of Mo
Single column ion-exchange (Pearce et al. 2009)	AG1X8	2	1 M HF/0.5 M HCl	4	Condition
			1 M HF/0.5 M HCl	6–10	Load
			1 M HF/0.5 M HCl	10	Matrix removal
			4 M HCl	8	Matrix removal
			3 M HNO ₃	12	Collection of Mo

537

538

539 Table 3. MC–ICP–MS operating conditions and Faraday collector configurations.

Neptune									
RF power	1165 W								
Ar cooling gas flow rate	16.00 L min ⁻¹								
Ar auxiliary gas flow rate	0.93 L min ⁻¹								
Ar sample gas flow rate	0.91 L min ⁻¹								
Sample cone	Ni sampler cone								
Skimmer cone	Ni X-skimmer cone								
Instrument resolution	400 (low)								
Integration time	4.194 s								
Idle time	3 s								
Cup configuration	L4	L3	L2	L1	C	H1	H2	H3	H4
	⁹¹ Zr	⁹² Mo	⁹⁴ Mo	⁹⁵ Mo	⁹⁶ Mo	⁹⁷ Mo	⁹⁸ Mo	⁹⁹ Ru	¹⁰⁰ Mo
Aridus II									
Sweep gas flow rate	3.45 L min ⁻¹								
Nitrogen gas	0.05 L min ⁻¹								
Spray chamber temperature	110 °C								
Desolvator temperature	160 °C								

540

541 Table 4. $\delta^{98/95}\text{Mo}$ values and Mo concentrations obtained for the BPHA column
 542 fractionation and Mo yield tests.

		Unspiked solution	Spiked solution
$\delta^{98/95}\text{Mo}_{\text{SRM 3134}} (\text{‰})$	solution 1	0.00 ± 0.07	0.04 ± 0.08
	solution 2	0.03 ± 0.08	-0.01 ± 0.06
	solution 3	0.02 ± 0.07	-0.01 ± 0.05
	Mean (2s)	0.02 ± 0.03	0.01 ± 0.06
Mo concentration (ug g ⁻¹)	solution 1	0.5402	0.5419
	solution 1	0.5342	0.5389
	solution 1	0.5365	0.5385
	Mean	0.5369	0.5398
	Stock solution	0.5429	0.5429

543

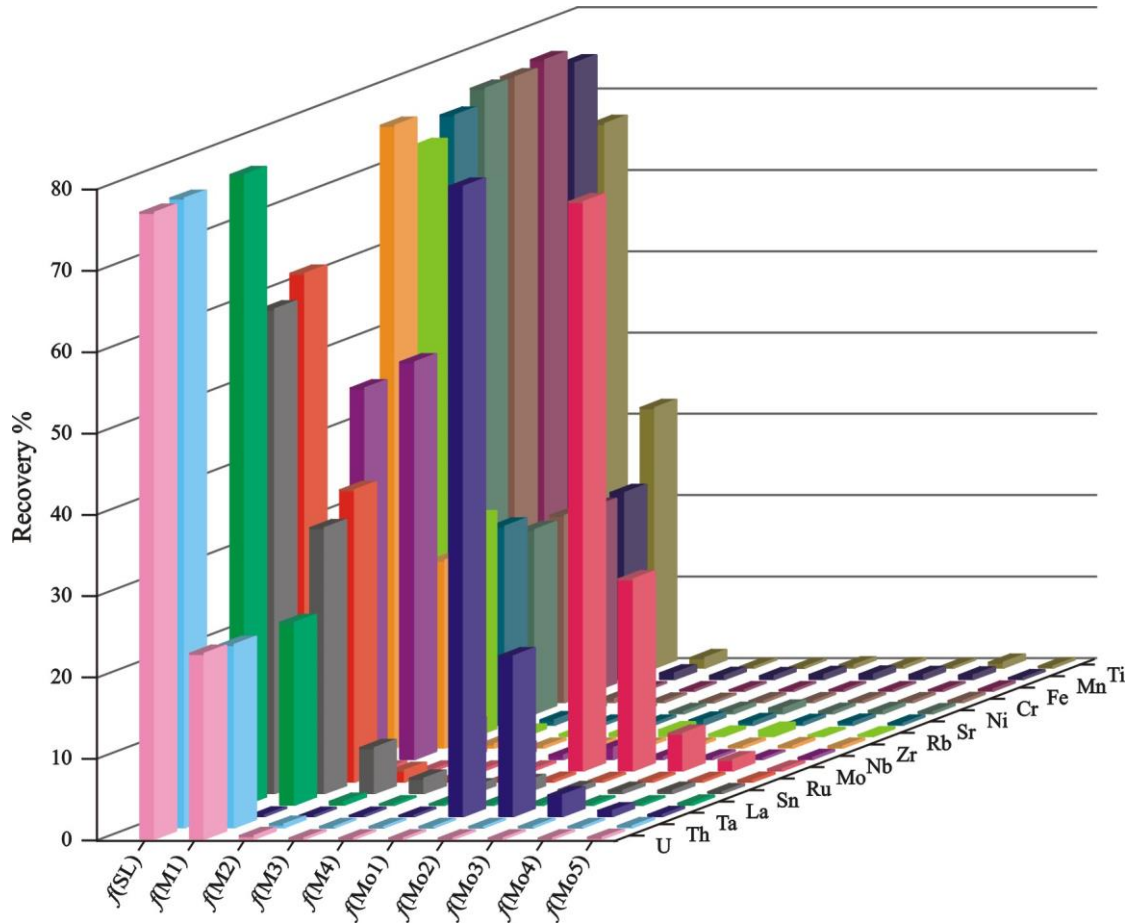
544 Table 5. $\delta^{98/95}\text{Mo}_{\text{SRM 3134}}$ and Mo concentration data determined for different
 545 geological materials.

Sample	$\delta^{98/95}\text{Mo}_{\text{SRM3134}}$ (%)	Mo concentration ($\mu\text{g g}^{-1}$)
BHVO-2	-0.11 ± 0.08	2.84
BHVO-2	-0.05 ± 0.08	3.79
BHVO-2	0.01 ± 0.06	5.15
Average (2 sd)	-0.05 ± 0.11	3.9 ± 2.3
Pearce et al.(2009)	0.08 ± 0.10	4.48
GBW07316	-0.56 ± 0.05	5.55
GBW07316	-0.58 ± 0.10	5.14
GBW07316	-0.66 ± 0.05	5.27
Average (2 sd)	-0.60 ± 0.10	5.32 ± 0.42
Previously reported [8,10]	-1.0 to 1.6	3.4 to 132
Reference value		5.70
IAPSO	2.01 ± 0.09	0.010
IAPSO	2.01 ± 0.10	0.010
IAPSO	1.98 ± 0.11	0.010
Average (2 sd)	2.00 ± 0.03	0.010
Greber et al.(2012)	2.08 ± 0.09	0.010

546

547 **Figure captions**

548 **Fig.1** Elution profile for a synthetic multi-element solution from a BPHA resin
549 column packed with 0.5 mL resin. $f(\text{SL}) = 2 \text{ mL of } 0.1 \text{ mol L}^{-1} \text{ HF}/1 \text{ mol L}^{-1} \text{ HCl}$ for
550 sample loading, $f(\text{mx}) = 2 \text{ mL of } 0.1 \text{ mol L}^{-1} \text{ HF}/1 \text{ mol L}^{-1} \text{ HCl}$ for matrix elution,
551 and $f(\text{Mox}) = 2 \text{ mL of } 6 \text{ mol L}^{-1} \text{ HF}/1 \text{ mol L}^{-1} \text{ HCl}$ for Mo elution.

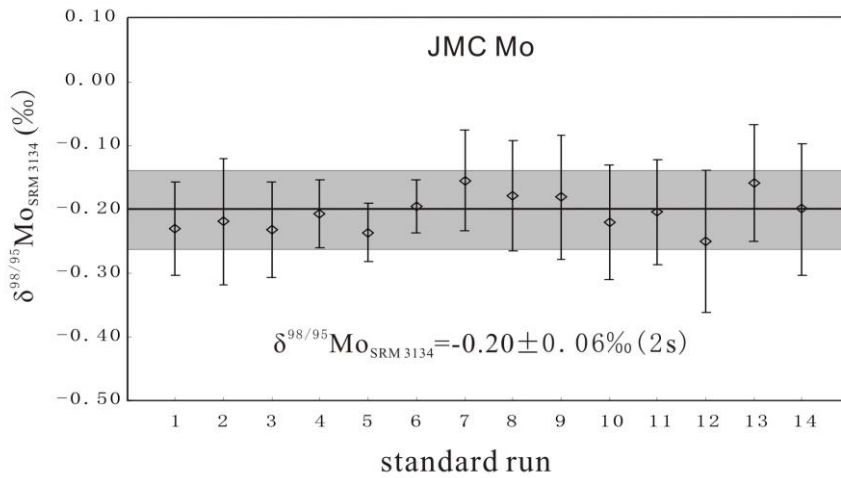
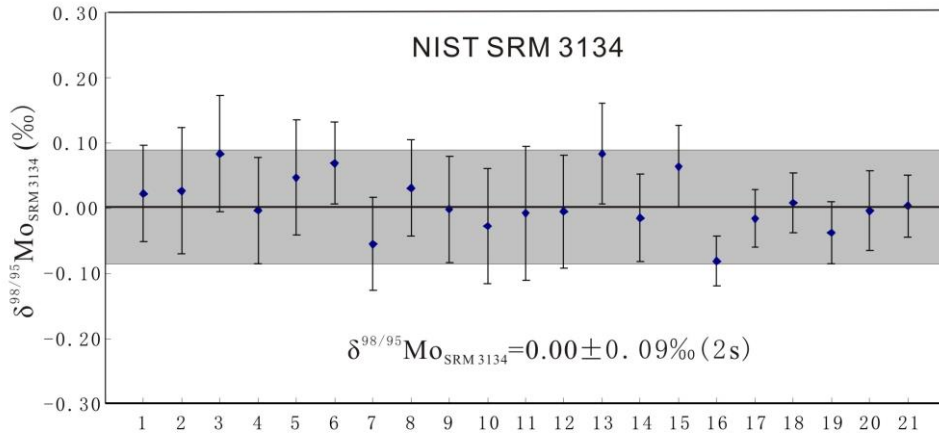


552

553

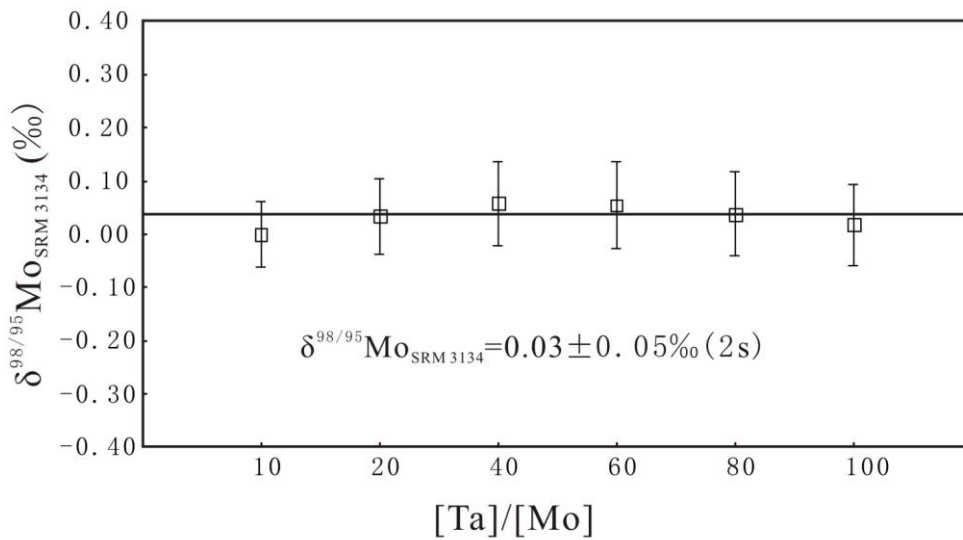
554

555 **Fig.2** Plot of $\delta^{98/95}\text{Mo}$ standard measurements for one session. The $\delta^{98/95}\text{Mo}$ values
 556 were normalized to the mean standard value of the session. Error bars are the internal
 557 uncertainty (2 se; standard error) and the external precision is represented by the gray
 558 band (2 sd; standard deviation).



559

560 **Fig.3** Effect of Ta doping on measured $\delta^{98/95}\text{Mo}$ values.



561