

GRACE hydrological monitoring of Australia: current limitations and future prospects

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ABSTRACT

The Gravity Recovery and Climate Experiment (GRACE) twin-satellite gravimetry mission has been monitoring time-varying changes of the Earth's gravitational field on a near-global scale since 2002. One of the environmentally important signals to be detected is temporal variations induced by changes in the distribution of terrestrial water storage (i.e., hydrology). Since water is one of Australia's precious resources, it is logical to monitor its distribution, and GRACE offers one such opportunity. We analyse the second and fourth releases (referred to as RL02 and RL04) of the 'standard' monthly GRACE solutions with respect to their annual mean. When compared to rainfall data over the same time period, GRACE is shown to detect hydrological signals over Australia, with the RL04 data showing better results. However, the relatively small hydrological signal typical for much of Australia is obscured by deficiencies in the standard GRACE data processing and filtering methods. Spectral leakage of oceanic mass changes also still contaminates the small hydrological signals typical over land. We therefore recommend that Australia-focussed reprocessing of GRACE data is needed for useful hydrological signals to be extracted. Naturally, this will have to be verified by independent in situ external sources such as rainfall, soil moisture and groundwater borehole piezometer data over Australia.

INTRODUCTION AND BACKGROUND

The Gravity Recovery and Climate Experiment (GRACE) is a joint US-German satellite mission dedicated to monitoring temporal and spatial variations in the Earth's external gravitational field on a near-global scale (e.g., Tapley et al., 2004a; 2004b; Wahr et al., 2004; Woodworth and Gregory, 2003; Rummel et al., 2002; Visser, 1999). This is done in a relatively cost-effective way by using two low-Earth orbiting (LEO) satellites operating in tandem (i.e., one following the other), coupled with precise orbit determination from a combination of the Global Positioning System (GPS) and inter-satellite K-band ranging.

As evidenced by much of the previous physical-geodetic research (e.g., Chao, 2003), the Earth's gravitational field has historically been treated as a largely static phenomenon, neglecting the albeit smaller temporal variations. For instance, AUSGeoid98 is assumed to be static, thus representing the gravity field as an average over the time the data used for its creation were collected (Featherstone et al., 2001).

Since the launch of the GRACE mission on 17 March 2002, it has become possible – for the first time – to monitor temporal variations in the Earth's external gravitational field on a near-global scale (e.g., Wahr et al., 2006; Riegger and Günter, 2004). GRACE allows these time-variations to be mapped every month (Luthcke et al., 2006) or less (e.g., Hu et al., 2006). While many factors cause the Earth's gravitational field to change with time, with tides being the most noticeable, another major factor is temporal variations in total terrestrial water storage, i.e., hydrology, on monthly timescales at spatial scales of the order of 1,000,000 km² (cf. Wahr et al., 1998).

Chen et al. (2006a) state that the GRACE mission duration has now been extended to 2010, extending its original five-year design life. In addition, plans are underway to launch a GRACE follow-on mission (R.S. Nerem, 2006 pers comm.), which will probably use lasers instead of K-band ranging, and thus improve the inter-satellite measurement accuracy from millimetres to microns, but the spatial resolution will probably not be improved much.

Numerous studies have demonstrated that GRACE can detect hydrological variations in major watersheds with active groundwater systems. Rodell and Famiglietti (1999) characterise water storage changes in 20 drainage basins of sizes varying from 130,000 km² to 5,782,000 km² to assess *detectability* of hydrological signals with respect to temporal and spatial variations. Tapley (2004a, 2004b) provide early results of the real application of GRACE satellites to detect hydrological signals in the Amazon-Orinoco basin. Following the results of Tapley (*ibid.*), several other authors have subsequently applied GRACE to detect hydrological signals in various situations and locations.

For instance, Ramillien et al. (2004, 2005) and Andersen et al. (2005) consider global time variations of hydrological signals and

variations of land-water storage from an inversion of two years of GRACE-derived geoids. Syed et al. (2005) examine total basin discharge for the Amazon-Orinoco and Mississippi river basins from GRACE and land-atmosphere water balance. Rodell et al. (2006) estimate groundwater storage changes in the Mississippi basin. Crowley et al. (2006) estimate hydrological signals in the Congo basin, while Schmidt et al. (2006b) and Swenson et al. (2003) use GRACE to observe changes in continental water storage. Awange et al. (2007) use GRACE to study the fall of Lake Victoria's water level in Africa.

In Australia, Rodell and Famiglietti (1999) assess the *detectability* of variations in continental water storage from GRACE in the Murray-Darling basin. Ellet et al. (2005) explore the potential of GRACE to assess large-scale hydrological models based on the simulation of Murray-Darling basin water storage for the period 2002-2003. Ellet et al. (2006) subsequently provide a framework for assessing the potential of GRACE to provide a new insight to the hydrology of the Murray-Darling basin.

Further examples of GRACE application to monitoring terrestrial hydrological variations include, e.g., Seo and Wilson (2005), Seo et al. (2006), Winsemius et al. (2006a; 2006b), Boy and Chao (2002), Hinderer et al. (2006), Llubes et al. (2004), Swenson et al. (2003), and Swenson and Milly (2006). Clearly, it is an active area of study.

However, with the exceptions of Rodell and Famiglietti (1999) and Ellett et al. (2005, 2006), the above-cited studies have been conducted where hydrological signals (actually gravitational field signals due to changes in the mass of terrestrial water) are large, which is not the case over most of Australia, especially in the current times of drought. Also, relatively less work has been done at smaller spatial scales (e.g., Swenson et al., 2006; Rodell and Famiglietti, 2001, 2005), principally due to the limited resolution of GRACE (because of the decrease in gravitational strength at the satellites' altitude), coupled with deficiencies in the data processing methods (discussed later).

This paper examines the current limitations of using GRACE to detect terrestrial water-mass variations over Australia from the 'standard' monthly gravity solutions (cf. Luthcke et al. 2006a). By 'standard', we mean that we use the fully normalised spherical harmonic coefficients as provided by one GRACE-processing group, rather than processing the raw data ourselves. The former has proven to be particularly challenging because errors in the high-frequency (i.e., smaller scale) data from GRACE collude with the relatively small hydrological signal typical for much of Australia, making the hydrological component harder to detect from space.

However, based on many studies reported elsewhere (e.g., Beylkin and Cramer, 2004; Fengler et al., 2006; Han et al., 2005a; Lemoine et al., 2007; Schmidt et al., 2006a, 2006b, 2007; Rowlands et al., 2005; King et al., 2006; Bauer et al., 2007; Hu et al., 2006; Han et al., 2005c), there might be some scope for improved data processing and

filtering techniques to be applied in an Australia-specific way to profit as much as possible from GRACE and its follow-on. By Australia-specific, we mean that existing techniques should be adapted for the Australian situation, or new techniques that suit Australia be devised.

Essentially, we have now reached an era of *environmental geodesy*, where high-precision modern geodetic observations are helping scientists from many other disciplines to get a better understanding of the physics behind environmental change (e.g., Leuliette et al., 2002; Vespe and Rutigliano, 2005; Davis et al., 2004). A key example is monitoring contemporary deglaciation through observed changes in external gravitation (e.g., Baur et al., 2007; Kuhn et al. 2006; Frappart et al. 2006; Velicogna and Wahr 2005; Velicogna et al. 2005; Chen et al. 2006a, 2006b). The physical-geodetic contribution to hydrology is yet another example, shown here.

KNOWN LIMITATIONS OF GRACE

In order to detect temporal gravitational variations at smaller spatial scales, the satellite(s) being tracked should be in as-low-as-possible orbits (close to the mass-change sources), with the proof masses in them being observed isolated, as-best-as-possible, from the perturbing effects of atmospheric drag.

LEO satellites have to be used for satellite gravimetry because of Newton's inverse square law of decay of the Earth's gravitational field with distance. The GRACE mission's altitude is 450 km, which permits the reliable detection of gravity field changes to spherical harmonic degree ~60 (Koch, 2005), corresponding to a spatial resolution of ~660 km (or ~330 km half-wavelength) at the equator.

Large time-varying gravitational signals, notably variations due to atmosphere- and ocean-mass tides, have to be removed during the processing of GRACE data so that the mission reveals time variations of hydrological, or other, interest. These de-aliasing models (for correcting short term; e.g., six-hour for tidal variations) have to be used to mitigate the propagation of spurious signals into the GRACE solutions (e.g., Han et al., 2004; Flechtner et al., 2006; Schrama and Visser, 2007; Knudsen and Andersen, 2002). However, these de-aliasing models are not perfect, so some spectral leakage and aliasing still remains in the GRACE solutions.

Probably the largest problem in GRACE solutions is that some of them are plagued with striping, where dominantly north-south strips associated with the GRACE satellite ground tracks tend to dominate the high-frequency component of the GRACE-derived gravitational signal (see, e.g., Fig. 9 in Tapley et al., 2005). There remains some conjecture as to the exact cause of the striping, but it is thought to be mostly due to too much weight being placed on the along-track K-band ranging data (e.g., Swenson and Wahr 2006; Schrama and Visser 2007), coupled with inaccurate de-aliasing models and the mission configuration.

Other GRACE errors relate to spectral leakage from time-varying ocean-based masses, which propagate spatially to mask any weak hydrological signals of interest (shown later). Finally, Chen et al. (2006c), among many others, point out that the high-degree and -order spherical harmonics of GRACE time-variable gravitational field models are dominated by noise.

CURRENT GRACE CONTRIBUTION TO AUSTRALIAN HYDROLOGY

The application of GRACE to monitoring hydrological variations works well (i.e., in terms of spatial resolution of the signal that can be detected) for a global model (e.g., Wu et al., 2006; Andersen and Hinderer, 2005) or for very large scale/volume drainage basins such as the Amazon, Mississippi, Ganges and Zambezi. Rodell and Famiglietti (2001) recommend basin sizes greater than 200,000 km² (e.g. ~450 km x ~450 km) for a precision of a few mm in equivalent water thickness (EWT) derived from GRACE. In terms of GRACE hydrology, the gravitational field data are routinely converted into units of mm of EWT using the algorithms in Wahr et al. (1998).

More recent studies (e.g., Smith et al. 2005) indicate that an area of at least 500,000 km² (e.g. ~700 km x ~700 km) and a time-period of one month are required for better EWT estimates. Areas of smaller spatial extent suffer from increased noise in high degrees and orders and spectral leakage of the gravitational effects from surrounding masses (Han et al. 2006), notably the oceans. Therefore, for efficient monitoring of Australia's stored water, two key issues emerge:

1. A desire to have a higher temporal resolution than the 30-day 'standard' GRACE data processing epochs (cf. Luthcke et al. 2006). If the variation in water-mass can be known sub-monthly rather than monthly, this would greatly enhance water management issues in Australia (cf. Hu et al. 2006).
2. A desire to exploit GRACE data to provide hydrological variations at higher spatial resolutions so as to cater for smaller basins and freshwater aquifers of less than 200,000 km² in area, which is the case for most of Australia.

Whereas GRACE-monitoring of stored water has been successfully tested for and applied to very large/volume drainage basins, as discussed earlier, the Australian case presents a more complicated scenario due, largely, to spectral-leakage contamination from oceanic mass-change signals coupled with the relatively small hydrological signal that is typical over most of Australia. As will be shown later, spectral leakage from the oceans obscures the land hydrological signal, sometimes for a considerable distance.

The current Australian drought has both positive and negative effects on the application of GRACE to Australian hydrology. On the positive side, it provides the opportunity for the application of GRACE to measure how Australia has dried out in the last five years and study whether rainfall actually replenishes the water supply. On the negative

side, the hydrological signals may have become too weak to be detected using 'standard' GRACE processing methods. Nevertheless, GRACE could be used to study variations in stored water over Australia, provided that data errors are reduced through proper filtering and data processing techniques.

Finally, the 'standard' monthly GRACE gravity field solutions are expressed in terms of fully normalised spherical harmonics, which are globally supported basis functions (cf. Blais and Provins, 2002). Spherical harmonics are the standard for representing the global external gravitational field. However, they are oscillatory functions that depend on superposition cancellation and addition to represent local features. In principle, there is a global mapping between every observation and every coefficient (e.g., Moritz, 1980; Colombo, 1989; Pavlis et al., 1996). This makes it difficult, if not impossible, to adjust the spatial frequency contents locally (i.e., using global spherical harmonic models for local mapping). While Chen et al. (2007) have used basin functions to try to mitigate this problem, they acknowledge that it is imperfect and still permits spectral leakage.

As such, it is also important to seek alternative representations of the Earth's external gravitational field that allow localised studies to be conducted from GRACE, and that they are not contaminated by data or mis-modelling errors from other areas. Some potential candidates will be discussed later.

Australia-wide hydrology

To indicate the capability of GRACE to monitor stored water over Australia, two data sets, namely RLO2 (released in 2006) and RL04 (released in May 2007), will be used to compute hydrological signals (EWT) over Australia for the year 2004 as an example. The purpose of using these two releases is to assess any improvement in release RL04 compared to RL02 in the Australian context. We use the 'standard' monthly GRACE solutions published by the Center for Space Research (CSR) at the University of Texas at Austin, USA, (<http://www.csr.utexas.edu/grace/>) to generate EWT for Australia from the Wahr et al. (1998) algorithm.

We acknowledge that other GRACE solutions are published by other groups; these being GFZ (GeoForschungZentrum, Germany), JPL (Jet Propulsion Laboratory, USA) and GRGS (Groupe de Recherche de Géodésie Spatiale, France). We chose the RL02 and RL04 CSR solutions mainly for reasons of convenience, but also to exemplify some of the problems with GRACE-based hydrology in Australia. As such, the comments and approaches in this paper apply to the CSR solutions only, and not all other GRACE solutions from other groups.

The monthly spherical harmonic coefficients are computed relative to the annual mean (i.e., $\text{coeff} = \text{coeff}(\text{month}) - \text{coeff}(\text{annual mean})$) for each month during 2004 as a typical example. The correlated-error filter proposed by Swenson and Wahr (2006) was applied to the

'standard' monthly spherical harmonic coefficients from CSR, followed by smoothing using a Gaussian filter of radius 500km. The filtering and smoothing reduce the south-north striping effects, which had previously made the detection of hydrological signals over Australia extremely difficult from RL02 of the CSR solutions. Recall that the need to apply Gaussian filtering is not always necessary for other GRACE models (cf. Lemoine et al., 2007).

Figures 1 and 2 show EWT computed from RL02 and RL04 respectively. Although most months show some similarities, August and September in 2004 show RL04 to be smoother than RL02, indicating the improved data processing techniques used in RL04. The most discernible EWT signals occur in northern Australia during February, March and April 2004, which are sub-tropical regions that often experience heavy rainfall during these months.

Comparison between GRACE signals (Fig. 2) and rainfall data for the same months in 2004 (Fig. 3) shows quite a good spatial correlation between the high rainfall in February (more than 300 mm) with the GRACE hydrological signals for the same month. The monthly rainfall data were provided by the Australian Bureau of Meteorology (cf. Jones and Weymouth, 1997), and have been generated using Barnes's (1994a,b,c) successive correction technique, which applies a weighted average to data reported within set grids across Australia (see, e.g., <http://www.bom.gov.au/climate/austmaps/mapinfo.shtml#variables>).

In order to be more consistent with the GRACE data presentation (Figures 1 and 2), the monthly rainfall variations in Figure 3 are computed relative to the annual mean (i.e., monthly total minus annual mean) for each month during 2004. It is important to acknowledge that a one-to-one correlation can never be expected from such a comparison (i.e., GRACE versus rainfall) because of many other factors that affect the mass-transport of terrestrial water. For instance, evapotranspiration and the rates of rainfall run-off will affect their temporal and spatial correlation. Nevertheless, we offer a first-look analysis of what may be the major signals.

In Figure 3, May 2004 shows hydrological signals in central Australia. These signals are detected by the GRACE satellites in the subsequent two months of June and July. From June to August, the Murray-Darling basin contains signals that are detected by GRACE in August and September (Figure 2). These examples indicate that the rainfall water has been retained for two months following the rains before being detected by GRACE. Such information may be relevant for water resource management and policy makers.

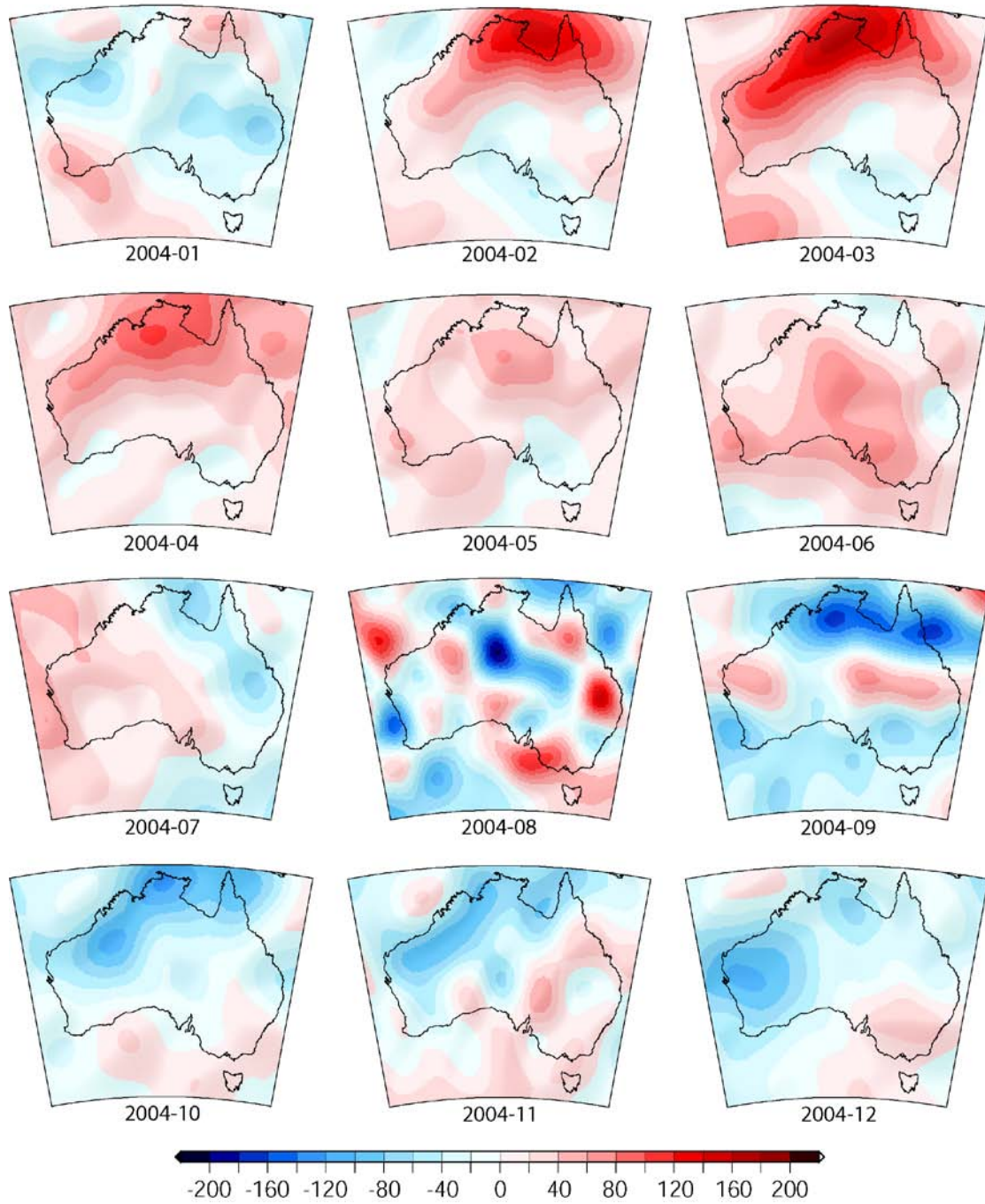


Figure 1. Monthly variations (in mm) of EWT from the mean over Australia in 2004 from CSR's GRACE release RL02 data (Lambert projection).

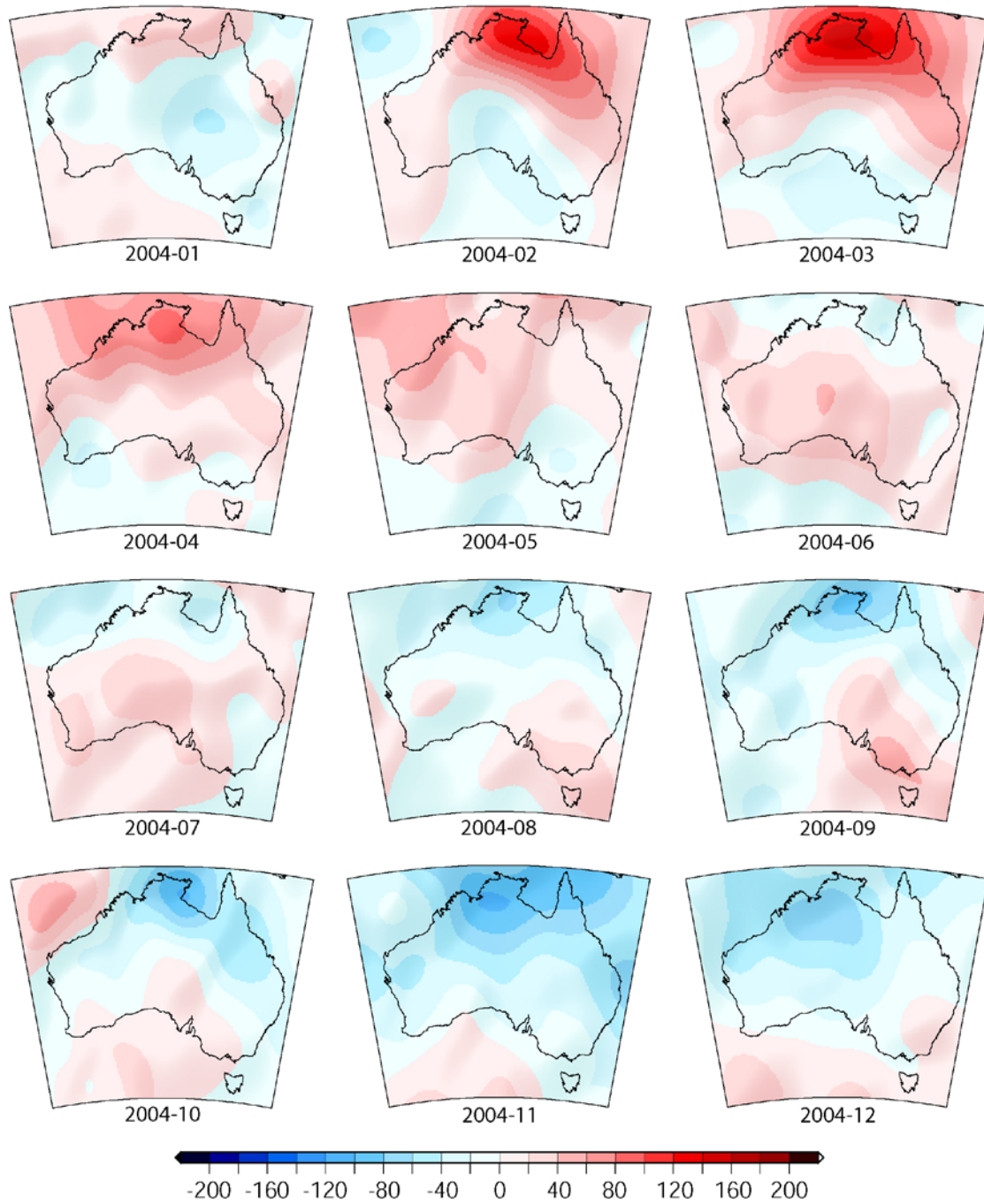


Figure 2. Monthly variations (in mm) of EWT from the mean over Australia in 2004 from CSR's GRACE release RL04 data (Lambert projection).

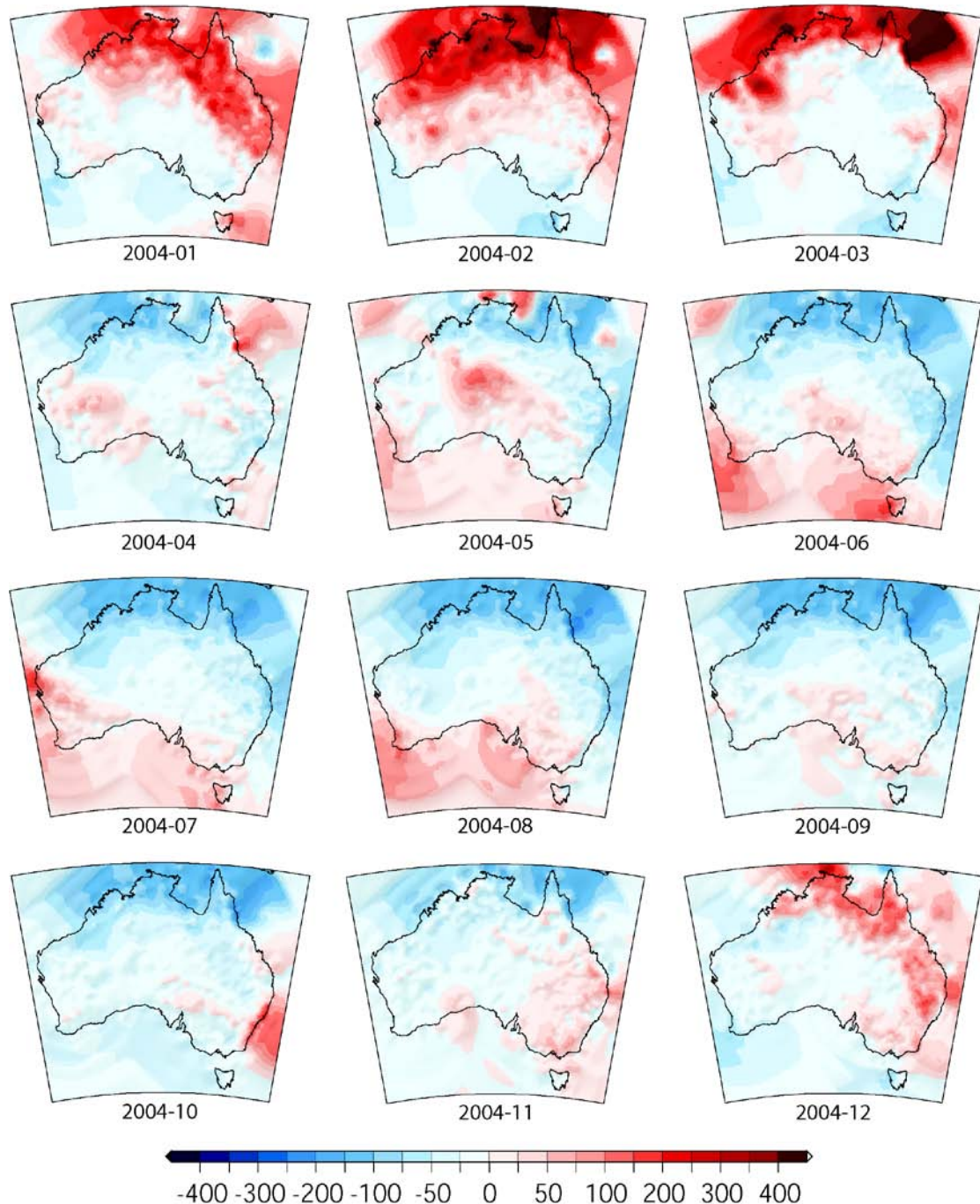


Figure 3. Monthly rainfall variations (in mm) from the mean over Australia in 2004 from Bureau of Meteorology rainfall data (Lambert projection)

While the signals discussed above are *detectable* by GRACE, weaker signals (e.g., April in Figure 3) are masked/obscured by spectral leakage, even in the RL04 data (Figure 2). This spectral leakage from ocean water-mass changes are dominant in April, May, June, November and December (Figure 2). This shows that ocean water-mass changes have not been modelled correctly, especially in the geodetically problematic coastal zone (cf. Deng and Featherstone, 2006). Therefore, improved filtering techniques are needed to remove aliasing (Han et al., 2004; Flechtner et al., 2006; Thompson et al., 2004),

and spectral leakage from the oceans (cf. Chen et al., 2006c).

DISCUSSION

The first-look results here show the potential for Australian hydrology to be monitored using GRACE, but only now for the improved CSR RL04 data that have been filtered and Gaussian-smoothed. The fact that some of the monsoonal seasons in the north of Australia appear in the GRACE data indicate that temporal and spatial variations in terrestrial water can indeed be detected from space. Also, the Murray-Darling basin has an area of 1,061,469 km², which is much greater than the 200,000 km² to 500,000 km² areas used in other studies elsewhere. However, hydrological signals likely in the Murray-Darling basin are much smaller than in other major drainage basins around the world (e.g., Amazon, Mississippi, Ganges and Zambezi). Nevertheless, there appears to be some modest amount of variability in and around the Murray-Darling basin detected from GRACE, which is encouraging (e.g., Figure 3 for September).

As opposed to large basins, significant water 'users' in Australia, e.g., farmers in the south-western Australian wheat-belt region or managers of freshwater aquifers all over Australia, are usually interested in water variation at much higher spatial resolution (e.g., ~200 km x ~200 km) than can be resolved by the CSR 'standard' monthly GRACE solutions. However, given the ultimate limitation of GRACE and GRACE follow-on missions imposed by atmospheric drag, satellite gravimetry will possibly never deliver detailed hydrological data at such small scales. Nevertheless, it is important to 'push the boundaries' to gain the most information from satellite-based hydrology.

In order to obtain any higher spatial resolution, more robust mathematical approaches that are applicable to small-scale hydrological mapping need to be investigated. Several procedures have been proposed and tested in other areas that could possibly be used in Australia to achieve higher spatial (less than 200,000km²) and temporal (sub-monthly) resolution hydrological variations from GRACE. Options include the wavelet *approach* (Beylkin and Cramer 2004; Fengler et al. 2006; Schmidt et al. 2006a), regional inverse method (Han et al. 2005b) and mascons (Lemoine et al. 2007). Importantly, these methods avoid the spherical harmonic approach, thus allowing better space-localisation. Indeed, there may be other approaches currently being developed, or ones that we can devise here in Australia.

Therefore, there is a need to analyse more robust method(s) that will better exploit GRACE data and provide timely small-scale variations in stored water over a short time frame (cf. Rowlands et al. 2005).

CONCLUSIONS AND RECOMMENDATIONS

From the 'standard' monthly GRACE solutions from CSR, even for the more recent RL04 data release, there is considerable influence of interfering effects such as GRACE de-aliasing errors and spectral

leakage that we attribute to ocean-mass variations. These impede the reliable detection of the relatively small hydrological signals expected in Australia. This is especially the case because it is the driest continent, and has been experiencing serious drought in many regions during the lifetime of the GRACE mission.

There is therefore need to investigate and apply Australia-specific techniques to better-estimate small hydrological signals from GRACE. The fundamental tasks required that we recommend are:

- Analysis of wavelet, regional inversion and mascon methods to establish if any best fits Australia, coupled with a search for any new Australia-specific method;
- Development, refinement and in-situ testing of whichever method selected from above to improve its capability for Australia;
- Application of the new method(s) to reprocess GRACE data to provide timely variations in Australia's stored water at both small and large scales to within, say, 14 days or less.

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