35 Key words: Electrical Resistivity Tomography, Engineered Barrier System, monitoring, non-

36 intrusive, geological disposal.

37

38

Introduction

- 39 Deep geological repository is favoured by many countries as a technically feasible and safe
- 40 programme for long-term disposal of high-level radioactive waste (Bredehoeft et al. 1978).
- 41 Although the selected host rock varies from country to country, all programmes consider the
- 42 implementation of an Engineered Barrier Systems (EBS) to directly protect and isolate the
- 43 waste. The material selected for the buffer surrounding waste canister as well as the material
- 44 that will be used to seal off the disposal galleries from the shafts leading to the surface is
- 45 generally based on compacted bentonite or bentonite/sand mixtures (Sellin & Leupin 2014).
- The EBS is subjected to an inward water flow from the host rock and an outward heat flux
- 47 from the radioactive waste (Lin et al. 1995; Rothfuchs et al. 2004; Jockwer et al. 2006; White
- 48 et al. 2017). Monitoring changes in water content and temperature is therefore the key to
- 49 assess the performance of the EBS. EBS monitoring during the operational period cannot be
- achieved via wired sensors installed in the buffer because wires can provide a preferential
- pathway for radionuclide leakage as well as for water (White et al. 2017).
- 52 Geophysical electrical monitoring is potentially an ideal technique for geophysical diffuse
- 53 monitoring of the EBS because (i) it can be designed in a non-intrusive fashion,(ii) it allows
- capturing local anomalies that local sensors cannot spot, and (iii) electrical resistivity is very
- sensitive to changes in water content and temperature and is therefore very convenient to
- monitor the EBS (Danielsen & Dahlin 2010; Korteland & Heimovaara 2015; Merritt et al.
- 57 2016; Carey et al. 2017; López-Sánchez et al. 2017; Wang et al. 2017; Cosenza et al. 2007;
- Hermans et al. 2015; Merritt et al. 2016; Carey et al. 2017; López-Sánchez et al. 2017; Wang
- 59 et al. 2017).
- 60 Electrical resistivity tomography (ERT) is a well-established geophysical technique that uses
- 61 injection of electrical currents and measurements of the resulting voltage differential at the
- 62 earth's surface or in boreholes. This generates pseudo-sections displaying apparent resistivity
- as a function of the location and electrode spacing, which in turn provides an initial picture of
- 64 the resistivity distribution. An inversion process of the measured data is necessary for the
- 65 final interpretation of the resistance data. This process transforms the apparent resistivity
- into 2D or 3D images of the bulk electrical resistivity of the subsurface model, which is
- discretised into a distinct number of elements of homogeneous resistivity.
- 68 ERT surveys have been routinely used in water exploration and contaminant flow detection
- 69 (de Lima et al. 1995; D. J. LaBrecque et al. 1996; Benson et al. 1997; Martinez-Pagan et al.
- 70 2009; Deceuster et al. 2013; Ntarlagiannis et al. 2016), engineering site investigations (Rucker
- 71 et al. 2009; Sentenac & Zielinski 2009; Banham & Pringle 2011; Jones et al. 2012, 2014),
- 72 location of buried artefacts or structures in archaeological surveys (Tonkov & Loke 2006;
- 73 Ullrich et al. 2007; Negri et al. 2008; Leucci & Greco 2012), as well as providing geological and

- 74 hydrogeological site information (Ganerød et al. 2006; Ramachandran et al. 2012; Aning et al.
- 75 2013).
- 76 ERT in boreholes has proven useful for environmental investigations (Daily & Owen 1991;
- 77 Daily et al. 1995; D. LaBrecque et al. 1996; French et al. 2002; Guérin 2005; Deceuster et al.
- 78 2006; Wilkinson et al. 2010). The method has also been demonstrated to be economically
- 79 efficient when using wells drilled for geotechnical pre-investigation tunnelling sites to obtain
- information about the geology between the wells (Denis et al. 2002). More recently,
- 81 investigations using ERT in borehole have been extended to a variety of other applications
- such as the characterization and monitoring of water infiltration (Oberdörster et al. 2010;
- 83 Coscia et al. 2011; Hermans et al. 2015) and monitoring CO₂ migration (Yang et al. 2015;
- 84 Schmidt-Hattenberger et al. 2016).
- 85 Previous researches conducted in repository-like conditions have demonstrated the potential
- of ERT in monitoring the EBS. Rothfuchs et al. (2004) could detect the water intake in an
- 87 experiment conducted in an area at the Aespoe Hard Rock Laboratory (HRL) in Sweden. ERT
- 88 electrode arrays were installed in the backfill, buffer and rock and the water saturation
- 89 changes in those three structures were monitored for a few years. Similarly, Furche & Scuster
- 90 (2014) have used ERT electrodes arrays installed in the Engineered Barrier Emplacement
- 91 Experiment in Opalinus Clay at the Mont Terri underground laboratory in Switzerland. Several
- 92 ERT surveys were conducted over the 11 years of operation of the experiment to monitor
- 93 water intakes in different areas of the experiment. However, in both these experiments, the
- 94 ERT electrodes were buried inside the EBS and this arrangement is not suitable for
- operational monitoring of the EBS. To the best of the authors' knowledge, there has been no
- attempt to date to investigate the use of the ERT technique in a non-intrusive fashion, i.e.
- 97 with the electrodes positioned outside the EBS.
- 98 This paper presents a mock-up scale test (ERT demonstrator) conceived within the EU project
- 99 'Modern2020' and implemented at the Underground Research Laboratory (URL) in
- Tournemire (France). It is intended to assess the capabilities of the Electrical Resistivity
- Tomography as a non-intrusive technique of monitoring the Engineered Barrier System under
- conditions as close as possible to the ones expected in the real repository. ERT electrodes
- were installed in two boreholes drilled at either side of the buffer to perform cross-borehole
- surveys. In the paper, three preliminary ERT surveys were carried out in January and
- November 2017 on the shaft before the emplacement of the bentonite. These surveys were
- aimed at a first assessment of the electrode installation technique, ERT measurement
- 107 protocols, and inversion procedures.

Description of Tournemire Underground Research Laboratory

110 Geological context

108

- 111 The French Institute of Radioprotection and Nuclear Safety (IRSN) uses Tournemire URL test
- site to conduct research on geological disposal of nuclear waste in clay formations (Cabrera

- et al. 2001; Gélis et al. 2010; Okay et al. 2013). Tournemire URL is located in southern France,
- in the western border of the Causses Basin (Cabrera et al. 2001; Okay et al. 2013).
- Fig. 1 shows the geological cross section of Tournemire. According to Okay et al. (2013) the
- intermediate formation, where the tunnel is located, correspond to marls and clay-rocks and
- is a good analogue of the Callovo-Oxfordian clay-rock in the Paris Basin, which is considered
- to be a potential host for the long-term storage of nuclear wastes in France.
- An old railway tunnel and six galleries are used to study the Toarcian formation (Fig. 2). In
- general, the Toarcian formation is mainly composed of illite (5–15% weight fraction),
- illite/smectite mixed-layer minerals (5–10% with a relative proportion of smectite of about
- 122 10%), chlorite (1–5%) and kaolinite (15–20%). This formation also contains 10–20% of quartz
- grains (weight fraction), 10–40% of carbonates (mainly composed of calcite with traces of
- dolomite and siderite) and 2–9% (in weight) of pyrite disseminated in the clay matrix
- spreading until 160 m deep from the tunnel (Cabrera et al. 2001; Okay et al. 2013).
- 126 The North-08 gallery
- 127 The area selected for the ERT demonstrator at the experimental site in Tournemire URL was
- the North-08 Gallery. The horseshoe cross-section of the North gallery is 3.7m tall and 4m
- wide along the floor. This gallery is 20m long oriented north-south (Fig. 2).
- On the left, approximately 3m of the area designated for the ERT demonstrator, there is a
- water infiltration experiment (WT-1) in progress, and on the right, approximately 5m of the
- 132 ERT demonstrator there is an empty borehole (GN1) of 0.1m in diameter and 7.15m long,
- located at 1.4m from the gallery floor.

Overview of the ERT demonstrator stages

- 136 The project was divided into three main stages. First stage, namely Stage 0, consisted of
- performing two blank tests before the installation of the EBS to establish the background
- resistivity of the rock mass. Blank test 1 comprised 2D surface measurements from the
- North-08 Gallery wall prior to the drilling of left and right boreholes and blank test 2
- constituted borehole measurements carried out from the left and right boreholes. Then, in
- 141 Stage 1, a shaft for the installation of the EBS was drilled and blank test 3 was carried out (Fig.
- 142 **3**).

- 143 The shaft is 60cm in diameter and approximately 9.05m long. The EBS is constituted by a 4m
- long mixture of bentonite pellets and powder, namely mixture 3, provided by NAGRA (Garitte
- et al. 2015). The average dry density of the pouring material is 1.45g/cm³ (Garitte et al. 2015).
- 146 Fig. 4 shows the particle size distribution of the material. The EBS will be closed off with a 2m
- long concrete plug. Hydration mats will be placed on both ends of the EBS and a heater on the
- bottom end. Two small access boreholes (Fig. 2) will be drilled perpendicular to the longitudinal
- 149 direction of the buffer to allow the installation of 16 local sensors: 8 Time Domain
- Reflectometry (TDR) and 8 temperature sensors, to measure water content and temperature
- as a way of cross-checking the geophysical measurements. For research purposes two lines of

- 152 16 electrodes each (0.24m spacing) will be buried inside the main shaft as well. The cross
- section of the EBS designed for this mock-up test and the instruments setup can be seen in Fig.
- 5. The installation of the EBS is scheduled to take place in July 2018 Error! Reference source not
- 155 **found.**.
- 156 The last stage, Stage 2, consists of regularly monitoring the changes in water content and
- temperature induced in the EBS using the local sensors and ERT measurements.
- 158 Several challenges surround this research experiment amongst them are: (1) electrodes
- 159 contact resistance problems (Day-Lewis et al. 2008; Danielsen & Dahlin 2010; Deceuster et al.
- 160 2013). The electrodes are installed in boreholes drilled in the rock. Usually, water is added
- within the borehole to ensure contact in these surveys. However, this resource is not an
- option for the ERT demonstrator since the electrode boreholes in question are horizontal. It
- is not possible to keep water in horizontal boreholes, thus continuous injection of water
- would be necessary in this situation, which would perturb the experiment; (2) data collection
- and processing (Oldenborger et al. 2005; Day-Lewis et al. 2008; Wilkinson et al. 2008;
- Deceuster et al. 2013). Borehole surveys involve several uncertainties, such as: position and
- alignment of electrodes, selection of the most appropriate arrays and measurements
- repeatability; (3) resolution and sensitivity of ERT in boreholes (D. LaBrecque et al. 1996;
- 169 Danielsen & Dahlin 2010; Tso et al. 2017).

Data collection of preliminary surveys

- 172 The main characteristics of the 2D ERT survey carried out during blank test 1 are presented in
- 173 Table 1. ARES II unit, manufactured by GF Instruments, was used for the data collection of
- 174 this blank test.

- 175 Two boreholes of 10cm in diameter and approximately 9.0m in length were drilled 1.20m
- apart, on either side of the position of the EBS, accommodating 32 electrodes spaced at
- 0.29m, within an inflatable PVC tube (Fig. 6), designed and manufactured by IRSN team. The
- inflatable system ensures contact between the electrodes and the borehole wall, as the
- injection of water into the boreholes would potentially disturb the resistivity of the study
- area hence it is out of question for this experiment. Cross-borehole measurements had been
- planned for blank test 2, however one of the connectors manufactured to enable the
- communication between the electrodes and ARES II unit did not work. As an alternative in-
- line borehole surveys (Fig. 7a) were performed in each borehole individually and the data
- 184 collected from both boreholes was combined. The multiplexer that accompanies this unit
- allows the connection of 48 electrodes in total (2 x 24 electrodes), hence the 8 most
- superficial electrodes in each borehole were not used in these measurements (Fig. 7a). Cross-
- borehole measurements were also performed using TERRAMETER LS ABEM unit including all
- 188 64 electrodes. For lack of familiarity with TERRAMETER LS ABEM unit at the time of blank test
- 2, the array used was a combination of AM-BN (Fig. 7b) where A and B are current
- electrodes and M and N are potential electrodes and AB-MN (Fig. 7c), that had been
- developed and implemented into the unit specifically for a previous IRSN research project.

- 192 Although part of the data collection of blank test 2 has been made on the boreholes
- independently using in-line borehole arrays, the data collected using ARES II and
- 194 TERRAMETER LS ABEM units have been processed together in cross-borehole format (values
- of geometric factor and hence resistivity were recalculated).
- 196 Prior to blank test 3, the shaft was drilled and two new sets of 32 electrodes each were
- designed, manufactured and installed into the boreholes by IRSN teamError! Reference
- 198 **source not found.** Cross borehole measurements were carried out using TERRAMETER LS
- 199 ABEM unit. The array used was AM-BN (Fig. 7b), based on experience gained from blank test
- 200 2 and recommendations of other researches (Day-Lewis et al. 2008; Wilkinson et al. 2008).

201

202

Results and discussions

- 203 Data quality
- 204 Contact resistance checks were carried out prior to the data collection of each survey. For the
- 205 2D surface survey, a paste of bentonite was used to coat the electrodes wherever needed to
- improve contact resistance. However, this resource could not be used for borehole surveys.
- 207 As suggested by Day-Lewis *et al.* (2008), cut-offs of $50k\Omega$ for borehole data and $20k\Omega$ for
- surface data were considered, since higher values may indicate that only a limited current
- 209 can be injected for that electrode pair. The largest contact resistance recorded for blank test
- 210 1 was 3.5 k Ω , i.e. all electrodes were included. The contact resistance collected before blank
- 211 tests 2 and 3 are plotted in Fig. 8. Some electrodes showed contact resistance larger than 50
- 212 $k\Omega$ and were discarded.
- 213 Both units used in the three blank tests offer stacking procedure. The stack procedure
- consists of collecting each quadripole several times and averaging the results. This procedure
- has two clear advantages: (1) random noise is averaged out, which improves signal-to-noise
- ratio and (2) the standard deviation (stacking error) provides means of quantifying error and
- defining data weights for inversion. For all blank tests carried out, the minimum number of
- 218 stacking selected was 4 and the maximum was 8. The maximum variation coefficient
- accepted was 2%. In practical terms, this means that if the average standard deviation of the
- first 4 measurements for a quadripole is greater than 2% then more measurements are going
- 221 to be collected for that quadripole up until the maximum number selected (equal to 8 in this
- case). The standard deviation of all data collected is then calculated and recorded, regardless
- of whether the value is higher or lower than 2%. Data with stacking errors larger than 3%
- were eliminated (Day-Lewis et al. 2008).
- The mean stacking error of blank test 1 was 0.16% and no recorded data had stacking errors
- larger than 3%. Fig. 9 illustrates the stacking error distribution of blank test 2 and blank test
- 3. The mean stacking error and the percentage of data larger than 3% obtained for each test
- carried out are detailed in Table 2. The lower stacking errors observed in blank test 1
- compared to the blank tests 2 and 3 can be justified by two main reasons, (i) the approaches
- used to improve the electrode contacts and (ii) the survey type. In blank test 1, where surface
- 231 surveys were carried out, bentonite was used to improve the contact between the electrode

- and the rock, while the electrode contacts of the other two blank tests, 2 and 3, were
- ensured only by pressure. In addition, the protocols used for blank test 1 were well-
- established 2D surface protocols with attested good sensitivities while the protocols of blank
- tests 2 and 3 had not been yet properly adapted.
- The length of the current pulse was selected equal to 300ms. Reciprocal measurements,
- 237 which involve swapping current and voltage electrode pairs, could not be collected due to
- 238 time constrains during the surveys.
- 239 Another concern for borehole surveys is the geometric factors, K. Geometric factors are
- 240 numerical multipliers used to convert the resistance *R* (voltage to current ratio) in apparent
- 241 resistivity ρ_a :
- 242 $\rho_a = K.R$
- The geometric factor depends on the geometry of each electrode spacing setup. For
- borehole surveys Wilkinson et al. (2008) demonstrated that large geometric sensitivities of an
- electrode configuration occur when the geometric factor, K, changes rapidly with position. In
- turn, this occurs when K is close to singular. In addition, K will also be large in the vicinities of
- the singularity. Due to several operational issues, the arrays used for data collection during
- blank test 2 were not the most suitable. Hence, a considerable amount of data collected
- presented large K values, therefore the data collected in blank test 2 were filtered based on
- 250 the geometric factor, i.e. data associated with geometric factors larger than 250m⁻¹ were
- discarded. Fig. 10 shows the distribution of apparent resistivity before and after filtering out
- measurements with high geometric factors for blank tests 2.
- Overall, contact resistance, stacking errors, and geometric factor errors were the three
- 254 features used to filter the data collected in the surveys performed for the ERT demonstrator.
- The percentage of total data removed from each survey is shown in Table 2.
- 256 Inversions
- To investigate the benefits of filtering data according to the strategies discussed in the
- 258 previous section, inversions were performed on both the original and filtered data sets for
- comparison. Table 2 shows the Root Mean Square (RMS) errors obtained from these
- inversions.
- 261 Inversions were performed using the commercially available software package Res2DInv®
- 262 (Loke 2015). After carefully testing numerous inversion settings (Day-Lewis et al. 2008), the
- default settings proved to be the most appropriate one. These settings were used for all
- 264 control parameters, which were kept identical for each inversion.
- 265 Tomograms plots generated from filtered data sets of blank test 1 2D surface survey
- 266 Schlumberger array; blank tests 2 in- and cross-hole array; and blank test 3 cross-hole
- array are shown in Fig. 11, Fig. 12 and Fig. 13 respectively. The geometric location of WT-1
- and GN1 are highlighted in the tomogram of blank test 1 (Fig. 11) as well as the future
- position of the main shaft and electrodes boreholes that at this stage had not yet been
- 270 drilled.

The result of blank test 1 presented in Fig. 11 shows that higher values of resistivity are found

at the surface. This is reasonable since the rock face exposed to the gallery presents lower

degree of saturation and, hence, higher values of resistivity. Below and around 0.5m the

resistivity of the rock mass is fairly homogeneous with values lower than 100Ω m, which is

consistent with the results shown in blank tests 2 (Fig. 12) and 3 (Fig. 13) and also with the

resistivity measured in the laboratory on core samples extracted from both boreholes

(average of 40Ω m). Cosenza et al. (2007) and Gélis et al. (2016) also reported similar results

in terms of resistivity of Tournemire's core samples and 2D ERT surveys in Tournemire URL

279 respectively.

276

284

287

296

303

307

In blank test 1, there is an area of high resistivity (between chainage 14 and 17m) that could

suggest the presence of an anomaly. This anomaly could be related to the WT-1 shaft, which

is empty in the first 3.4m. There is another area of high resistivity in the model between

chainage 12 and 13.2m that extends to almost 2m into the wall. From all the field data and

information gathered so and made available by the IRSN team, there is nothing in this latest

segment that could justify such a high resistivity. Thus, a possible interpretation of these

results is that the high resistivity along the segment 14 and 17m is an artefact and WT-1 shaft

is actually associated with the high resistivity area between 12 and 13.2m. To investigate the

issue further, an inversion was tested with a priori resistivity information of WT-1 and GN1.

The inversion results have created an even larger artefact of high resistivity over almost the

290 whole model and the RMS error of this inversion has doubled. As the RMS indicates the

291 mismatch between the forward and calculated models, these results were not considered

satisfactory. Therefore, it was speculated that the problem stemmed from a 2D inversion

algorithms used to invert data of 3D bodies located outside the image plane (Nimmer et al.

294 2008).

295 The empty shaft of WT-1 presents virtually infinite resistivity and is by-passed by the current,

which follows more conductive paths. The stainless steel lid (35cm thick) is located at 3.4m

depth into the WT-1 shaft likely affecting the resistivity measurements (although the lid itself

is outside of the area of the inversion). Furthermore, WT-1 is located towards the edge of the

area covered by the inversion model, which is highly affected by boundary effects. As a

result, WT-1 is not clearly detected.

301 Blank test 2 (Fig. 12) was a combination of data collected from arrays involving in-hole and

cross-hole quadripoles combinations. The data was processed in cross-borehole format,

treated according to the procedure described in the data quality session and inverted. Fig. 12

shows that the resistivity between the two boreholes is somehow homogeneous and lower

than 100Ω m. The area of higher resistivity around the electrodes and in the middle of the

model (around 5m depth) is most likely due to artefacts created by the noise survey. A

considerable number of negative apparent resistivity data was collected during blank test 2.

308 This negative apparent resistivity does not appear to be real, since virtually no negative

apparent resistivity remained after filtering the data according to the data quality procedure

310 (Fig. 10).

Blank test 3 (Fig. 13) has an empty shaft (0.6cm in diameter and 9.05m in length) in the 311 312 middle of the cross borehole model, which should be characterised by high resistivity values. 313 However, higher resistivity values (greater than 500 Ω m) can only be spotted in the first 2.0m of the model, close to the gallery wall. This inconsistency was expected due to the presence 314 315 of the shaft. The current flow is expected to act three dimensionally avoiding the volume of high resistivity. Inverting the data collected in the blank test 3 using a 3D algorithm would not 316 317 improve the results. The problem of this survey is the data collection itself. The main shaft 318 represents a 3D body characterised by virtually infinite resistivity. Although there has been a 319 significant improvement in the protocol used for blank test 3 when compared to the one 320 used on blank test 2, the site characteristics were very difficult to capture using 2D surveys. To test this hypothesis a 3D synthetic model was created reproducing the site characteristics 321 322 (Fig. 14a). The model has 2.4m x 2.6m x 10m with background resistivity of 40Ω m, replicating 323 the resistivity of the core rock samples tested in laboratory, and a shaft of $0.6 \text{m} \times 0.6 \text{m} \times 10^{-2}$ 324 9.05m in the middle with resistivity of 1E+15 Ω m, representing the empty shaft. The synthetic 325 data were created in 3D, without adding noise, but the protocol used was the same of blank test 3. Firstly, the data were inverted using a 3D algorithm (RES3DInv® - (Loke 2017)), and the 326 327 tomography result can be observed in Fig. 14b. Apart from a few artefacts of high resistivity 328 around the edges, the resistivity of the whole model is homogeneous and around $100\Omega m$. 329 Therefore, the high resistivity body representing the main shaft is not characterised in the 330 tomography results. Then, the same data were inverted using a 2D algorithm and the 331 tomography result is presented in Fig. 14c. The highest resistivity value observed is 250 Ω m in 332 the centre towards the bottom of the model. Outside this area, the resistivity of the model is 333 homogenous and around 100Ω m. The higher resistivity observed in the 2D inverted model is 334 not enough to characterise precisely the empty shaft. Therefore, the outcome shows that the 335 2D protocol used in blank test 3 was unable to capture the main empty shaft regardless of 336 the inverted algorithm used. 337 For the monitoring stages of this experiment, protocols need to be improved and tested by 338 means of forward modelling and sensitivity analysis to ensure the quality of the data 339 collected and consistency of the inversion results. The possibility of adding a third borehole 340 to install electrodes at the top of the main shaft is currently being examined. This additional 341 set of electrodes could improve the tomography images. In this way, the data can be

342343

344

Conclusions

This paper has presented the preliminary Electrical Resistivity Tomography surveys of the ERT demonstrator carried out in Tournemire URL. This demonstrator is aimed to investigate the potential of ERT as non-invasive monitoring of the thermo-hydraulic response of the Engineered Barrier System (EBS) during the operational stage. The blank test surveys have allowed characterising the resistivity of the host rock and, most importantly, have allowed identifying the most suitable ERT protocols to be adopted in the next stages of the project when the EBS will be put in place.

collected in a real 3D fashion and inverted using 3D algorithm.

- Results obtained from laboratory experiments performed on core samples extracted from
- different depths during the drilling process suggested that the resistivity of the host rock is
- homogeneous and around 40 Ω m. The homogeneously of the host rock was indeed
- confirmed by blank test 1 and blank test 2, with consistent resistivity values lower than
- 356 100Ω m. The methodology developed for the electrode installation based on the use of PVC
- half-tubes pushed against the borehole wall by inflatable pipes has proved to be successful.
- However, electrodes contact resistance remains a challenge that need to be addressed.
- Inspection of the tomograms derived from in- and cross-hole array has highlighted the
- drawbacks of the protocols used and suggested the modifications to be introduced in the
- next stage of the experimental programme. In particular, the lesson learned from the blank
- tests allowed the following actions to be put in place:
 - Since the problem is clearly 3D, electrodes should be placed in 3D configuration, i.e. a third electrode array should be added at the top of the main shaft to complement the two arrays located laterally to the main shaft (on the left-hand and right-hand sides respectively). In this way, data can be collected in 3D fashion and inverted using 3D inversion algorithms. This measure should reduce the appearance of artefacts and allow generating enhanced tomography images;
 - New measurement protocols suitable for in-hole and cross-hole need to be
 developed to allow for more efficient data collection in terms of measurement time
 and adequate geometric factors. To ensure the quality of the measurement protocols,
 sensitivity analysis should be carried out on various protocol datasets complemented
 by similar analysis using synthetic data via forward model;
 - Reciprocal data should be collected to allow for enhanced data quality control.

375

376

374

363364

365

366

367

368

369370

371

372373

Acknowledgements

- 377 The authors wish to acknowledge the support of the European Commission via the project
- 378 MODERN2020 'Development and Demonstration of monitoring strategies and technologies
- for geological disposal' (Grant Agreement number: 662177-Modern2020-NFRP-2014-2015)
- under the H2020 Euratom Research and Training Programme. We also thank ANDRA and
- 381 IRSN for the funding support. And we thank Patrice Desveaux and Bruno Combes for their
- support in the experiments and for manufacturing the electrodes.

383

384

References

- Aning, A.A., Tucholka, P. & Danuor, S.K. 2013. 2D Electrical Resistivity Tomography (ERT)
- Survey using the Multi-Electrode Gradient Array at the Bosumtwi Impact Crater ,. 3, 12–
- 387 27.
- Banham, S. & Pringle, J.K. 2011. Geophysical and intrusive site investigations to detect an
- abandoned coal-mine access shaft, Apedale, Staffordshire, UK. Near Surface Geophysics,
- **9**, 483–496.

- 391 Benson, A.K., Payne, K.L. & Stubben, M.A. 1997. Mapping groundwater contamination using
- 392 dc resistivity and VLF geophysical methods—A case study. Geophysics, 62, 80–86,
- 393 https://doi.org/10.1190/1.1444148.
- 394 Bredehoeft, J.D., England, A.W., Stewart, D.B., Trask, N.J. & Winograd, I.J. 1978. Geologic 395 Disposal of High-Level Radioactive Wastes- Earth-Science Perspectives.
- 396 Cabrera, J., Beaucaire, C., et al. 2001. Projet Tournemire – Synthèse Des Programmes de Recherche 1995–1999. Report #IPSN DPRE/SERGD. Paris. 397
- 398 Carey, A.M., Paige, G.B., Carr, B.J. & Dogan, M. 2017. Forward modeling to investigate 399 inversion artifacts resulting from time-lapse electrical resistivity tomography during 400 rainfall simulations. Journal of Applied Geophysics, 145, 39–49,
- 401 https://doi.org/10.1016/j.jappgeo.2017.08.002.
- 402 Coscia, I., Greenhalgh, S.A., et al. 2011. 3D crosshole ERT for aquifer characterization and 403 monitoring of infiltrating river water. Geophysics, 76, G49–G59,
- 404 https://doi.org/10.1190/1.3553003.
- Cosenza, P., Ghorbani, A., Florsch, N. & Revil, A. 2007. Effects of drying on the low-frequency 405 406 electrical properties of Tournemire argillites. Pure and Applied Geophysics, 164, 2043– 2066, https://doi.org/10.1007/s00024-007-0253-0. 407
- 408 Daily, W. & Owen, E. 1991. Cross-borehole resistivity tomography. Geophysics, 56, 1228-409 1235.
- 410 Daily, W., Ramirez, A., LaBrecque, D. & Barber, W. 1995. Electrical resistance tomography experiments at the Oregon Graduate Institute. Journal of Applied Geophysics, 33, 227-411 412 237, https://doi.org/10.1016/0926-9851(95)90043-8.
- 413 Danielsen, B.E. & Dahlin, T. 2010. Numerical modelling of resolution and sensitivity of ERT in 414 horizontal boreholes. Journal of Applied Geophysics, 70, 245–254, 415 https://doi.org/10.1016/j.jappgeo.2010.01.005.
- Day-Lewis, F.D., Johnson, C.D., Singha, K. & Lane Jr, J.W. 2008. Best Practices in Electrical 416 417 Resistivity Imaging: Data Collection and Processing, and Application to Data from 418 Corinna, Maine.
- 419 de Lima, O.A.L., Sato, H.K. & Porsani, M.J. 1995. Imaging industrial contaminant plumes with 420 resistivity techniques. Journal of Applied Geophysics, 34, 93–108,
- 421 https://doi.org/10.1016/0926-9851(95)00014-3.
- 422 Deceuster, J., Delgranche, J. & Kaufmann, O. 2006. 2D cross-borehole resistivity
- tomographies below foundations as a tool to design proper remedial actions in covered 423
- 424 karst. Journal of Applied Geophysics, 60, 68-86,
- 425 https://doi.org/10.1016/j.jappgeo.2005.12.005.
- 426 Deceuster, J., Kaufmann, O. & Camp, M. Van. 2013. Automated identification of changes in
- 427 electrode contact properties for long-term permanent ERT monitoring experiments.
- Geophysics, 78, E79–E94, https://doi.org/10.1190/GEO2012-0088.1. 428
- 429 Denis, A., Marache, A., Obellianne, T. & Breysse, D. 2002. Electrical resistivity borehole
- 430 measurements: Application to an urban tunnel site. Journal of Applied Geophysics, 50,

431	319-331, http	s://doi.org/1	0.1016/	'S0926-9851(02)00150-7.

- 432 French, H.K., Hardbattle, C., Binley, A., Winship, P. & Jakobsen, L. 2002. Monitoring snowmelt
- induced unsaturated flow and transport using electrical resistivity tomography. *Journal*
- 434 of Hydrology, **267**, 273–284, https://doi.org/10.1016/s0022-1694(02)00156-7.
- Furche, M. & Scuster, K. 2014. Long-Term Performance of Engineered Barrier Systems PEBS.
- 436 Ganerød, G.V., Rønning, J.S., Dalsegg, E., Elvebakk, H., Holmøy, K., Nilsen, B. & Braathen, A.
- 437 2006. Comparison of geophysical methods for sub-surface mapping of faults and
- fracture zones in a section of the Viggja road tunnel, Norway. Bulletin of Engineering
- 439 *Geology and the Environment*, **65**, 231–243, https://doi.org/10.1007/s10064-006-0041-
- 440 6.
- Garitte, B., Weber, H. & Müller, H.R. 2015. *Requirements, Manufacturing and QC of the Buffer Components Report LUCOEX WP2*.
- 443 Gélis, C., Revil, A., et al. 2010. Potential of electrical resistivity tomography to detect fault
- zones in limestone and argillaceous formations in the experimental platform of
- Tournemire, France. Pure and Applied Geophysics, 167, 1405–1418,
- 446 https://doi.org/10.1007/s00024-010-0097-x.
- 447 Gélis, C., Noble, M., Cabrera, J., Penz, S., Chauris, H. & Cushing, E.M. 2016. Ability of High-
- 448 Resolution Resistivity Tomography to Detect Fault and Fracture Zones: Application to
- the Tournemire Experimental Platform, France. Pure and Applied Geophysics, 173, 573–
- 450 589, https://doi.org/10.1007/s00024-015-1110-1.
- Guérin, R. 2005. Borehole and surface-based hydrogeophysics. *Hydrogeology Journal*, **13**,
- 452 251–254, https://doi.org/10.1007/s10040-004-0415-4.
- 453 Hermans, T., Wildemeersch, S., Jamin, P., Orban, P., Brouyère, S., Dassargues, A. & Nguyen, F.
- 454 2015. Quantitative temperature monitoring of a heat tracing experiment using cross-
- 455 borehole ERT. *Geothermics*, **53**, 14–26,
- 456 https://doi.org/10.1016/j.geothermics.2014.03.013.
- Jockwer, N., Wieczorek, K., Miehe, R. & Diaz, A.M.F. 2006. *Heater Test in the Opalinus Clay of the Mont Terri URL Gas Release and Water Redistribution*.
- Jones, G., Zielinski, M. & Sentenac, P. 2012. Mapping desiccation fissures using 3-D electrical
- resistivity tomography. *Journal of Applied Geophysics*, **84**, 39–51,
- 461 https://doi.org/10.1016/j.jappgeo.2012.06.002.
- Jones, G., Sentenac, P. & Zielinski, M. 2014. Desiccation cracking detection using 2-D and 3-D
- Electrical Resistivity Tomography: Validation on a flood embankment. *Journal of*
- 464 Applied Geophysics, **106**, 196–211, https://doi.org/10.1016/j.jappgeo.2014.04.018.
- 465 Korteland, S.A. & Heimovaara, T. 2015. Quantitative inverse modelling of a cylindrical object
- in the laboratory using ERT: An error analysis. *Journal of Applied Geophysics*, **114**, 101–
- 467 115, https://doi.org/10.1016/j.jappgeo.2014.10.026.
- 468 LaBrecque, D., Miletto, M., Daily, W., Ramirez, A. & Owen, E. 1996. The effects of noise on
- 469 Occam's inversion of resistivity tomography data. *Geophysics*, **61**, 538–548.

- LaBrecque, D.J., Ramirez, A.L., Daily, W.D., Binley, A.M. & Schima, S.A. 1996. ERT monitoring of environmental remediation processes. *Measurement Science and Technology*, **7**, 375–
- 472 383, https://doi.org/10.1088/0957-0233/7/3/019.
- Leucci, G. & Greco, F. 2012. 3D ERT Survey to Reconstruct Archaeological Features in the
- Subsoil of the 'Spirito Santo' Church Ruins at the Site of Occhiolà (Sicily, Italy).
- 475 *Archaeology*, **1**, 1–6, https://doi.org/10.5923/j.archaeology.20120101.01.
- Lin, W., Wilder, D.G., et al. 1995. A Heated Large Block Test for High Level Nuclear Waste

 Management 2. *In: 2nd International Conference on Mechanics of Jointed and Faulted*
- 478 Rock (MJFR-2). Vienna.
- Loke, M.H. 2015. RES2DINV. Rapid 2-D Resistivity & IP inversion using the least-squares method. 127P.
- Loke, M.H. 2017. Rapid 3-D Resistivity & IP Inversion Using the Least-Squares Method.
- 482 López-Sánchez, M., Mansilla-Plaza, L. & Sánchez-de-laOrden, M. 2017. Geometric factor and
- influence of sensors in the establishment of a resistivity-moisture relation in soil
- samples. *Journal of Applied Geophysics*, **145**, 1–11,
- 485 https://doi.org/10.1016/j.jappgeo.2017.07.011.
- 486 Martinez-Pagan, P., Faz, A. & Aracil, E. 2009. The use of 2D electrical tomography to assess
- pollution in slurry ponds of the Murcia region, SE Spain. Near Surface Geophysics, 7, 49–
- 488 61, https://doi.org/10.3997/1873-0604.2008033.
- 489 Merritt, A.J., Chambers, J.E., Wilkinson, P.B., West, L.J., Murphy, W., Gunn, D. & Uhlemann, S.
- 490 2016. Measurement and modelling of moisture-electrical resistivity relationship of fine-
- grained unsaturated soils and electrical anisotropy. Journal of Applied Geophysics, 124,
- 492 155–165, https://doi.org/10.1016/j.jappgeo.2015.11.005.
- 493 Negri, S., Leucci, G. & Mazzone, F. 2008. High resolution 3D ERT to help GPR data
- interpretation for researching archaeological items in a geologically complex subsurface.
- 495 *Journal of Applied Geophysics*, **65**, 111–120,
- 496 https://doi.org/10.1016/j.jappgeo.2008.06.004.
- Nimmer, R.E., Osiensky, J.L., Binley, A.M. & Williams, B.C. 2008. Three-dimensional effects
- 498 causing artifacts in two-dimensional, cross-borehole, electrical imaging. *Journal of*
- 499 *Hydrology*, **359**, 59–70, https://doi.org/10.1016/j.jhydrol.2008.06.022.
- Ntarlagiannis, D., Robinson, J., Soupios, P. & Slater, L. 2016. Field-scale electrical geophysics
- over an olive oil mill waste deposition site: Evaluating the information content of
- resistivity versus induced polarization (IP) images for delineating the spatial extent of
- organic contamination. *Journal of Applied Geophysics*, **135**, 418–426,
- 504 https://doi.org/10.1016/j.jappgeo.2016.01.017.
- Oberdörster, C., Vanderborght, J., Kemna, A. & Vereecken, H. 2010. Investigating Preferential
- Flow Processes in a Forest Soil Using Time Domain Reflectometry and Electrical
- Resistivity Tomography. *Vadose Zone Journal*, **9**, 350–361,
- 508 https://doi.org/10.2136/vzj2009.0073.
- Okay, G., Cosenza, P., Ghorbani, A., Camerlynck, C., Cabrera, J., Florsch, N. & Revil, A. 2013.
- Localization and characterization of cracks in clay-rocks using frequency and time-

511 512	domain induced polarization. <i>Geophysical Prospecting</i> , 61 , 134–152, https://doi.org/10.1111/j.1365-2478.2012.01054.x.
513 514 515	Oldenborger, G.A., Routh, P.S. & Knoll, M.D. 2005. Sensitivity of electrical resistivity tomography data to electrode position errors. <i>Geophys. J. Int.</i> , 1–9, https://doi.org/10.1111/j.1365-246X.2005.02714.x.
516 517 518 519	Ramachandran, K., Tapp, B., Rigsby, T. & Lewallen, E. 2012. Imaging of fault and fracture controls in the arbuckle-simpson aquifer, Southern Oklahoma, USA, through electrical resistivity sounding and tomography methods. <i>International Journal of Geophysics</i> , 1–10, https://doi.org/10.1155/2012/184836.
520 521	Rothfuchs, T., Miehe, R., Moog, H. & Wieczorek, K. 2004. <i>Geoelectric Investigation of Bentonite Barrier Saturation</i> .
522 523 524	Rucker, D.F., Levitt, M.T. & Greenwood, W.J. 2009. Three-dimensional electrical resistivity model of a nuclear waste disposal site. <i>Journal of Applied Geophysics</i> , 69 , 150–164, https://doi.org/10.1016/j.jappgeo.2009.09.001.
525 526 527 528	Schmidt-Hattenberger, C., Bergmann, P., Labitzke, T., Wagner, F. & Rippe, D. 2016. Permanent crosshole electrical resistivity tomography (ERT) as an established method for the long-term CO2monitoring at the Ketzin pilot site. <i>International Journal of Greenhouse Gas Control</i> , 52 , 432–448, https://doi.org/10.1016/j.ijggc.2016.07.024.
529 530 531	Sellin, P. & Leupin, O.X. 2014. The use of clay as an engineered barrier in radioactive-waste management - A review. <i>Clays and Clay Minerals</i> , 61 , 477–498, https://doi.org/10.1346/CCMN.2013.0610601.
532 533 534	Sentenac, P. & Zielinski, M. 2009. Clay fine fissuring monitoring using miniature geo-electrical resistivity arrays. <i>Environmental Earth Sciences</i> , 59 , 205–214, https://doi.org/10.1007/s12665-009-0017-5.
535 536 537	Tonkov, N. & Loke, M.H. 2006. A resistivity survey of a burial mound in the 'Valley of the Thracian Kings'. <i>Archaeological Prospection</i> , 13 , 129–136, https://doi.org/10.1002/arp.273.
538 539 540	Tso, C.H.M., Kuras, O., et al. 2017. Improved characterisation and modelling of measurement errors in electrical resistivity tomography (ERT) surveys. <i>Journal of Applied Geophysics</i> , 146 , 103–119, https://doi.org/10.1016/j.jappgeo.2017.09.009.
541 542	Ullrich, B., Guenther, T. & Ruecker, C. 2007. Electrical Resistivity Tomography Methods for Archaeological Prospection. <i>Geophysical Prospecting</i> , 1–7.
543 544 545	Wang, J., Zhang, X. & Du, L. 2017. A laboratory study of the correlation between the thermal conductivity and electrical resistivity of soil. <i>Journal of Applied Geophysics</i> , 145 , 12–16, https://doi.org/10.1016/j.jappgeo.2017.07.009.
546 547	White, M., Farrow, J. & Crawford, M. 2017. <i>Deliverable D2.1: Repository Monitoring Strategies and Screening Methodologies</i> .
548 549 550	Wilkinson, P.B., Chambers, J.E., Lelliott, M., Wealthall, G.P. & Ogilvy, R.D. 2008. Extreme sensitivity of crosshole electrical resistivity tomography measurements to geometric errors. <i>Geophysical Journal International</i> , 173 , 49–62, https://doi.org/10.1111/j.1365-

- 551 246X.2008.03725.x.
- Wilkinson, P.B., Meldrum, P.I., Kuras, O., Chambers, J.E., Holyoake, S.J. & Ogilvy, R.D. 2010.
- High-resolution Electrical Resistivity Tomography monitoring of a tracer test in a
- confined aguifer. *Journal of Applied Geophysics*, **70**, 268–276,
- https://doi.org/10.1016/j.jappgeo.2009.08.001.
- Yang, X., Lassen, R.N., Jensen, K.H. & Looms, M.C. 2015. Monitoring CO2migration in a
- shallow sand aquifer using 3D crosshole electrical resistivity tomography. *International*
- Journal of Greenhouse Gas Control, **42**, 534–544,
- 559 https://doi.org/10.1016/j.ijggc.2015.09.005.

- 561 Figure captions
- Fig. 1. Geological cross section of Tournemire URL.
- Fig. 2. Position of the galleries at the experimental site.
- Fig. 3. Overview of ERT demonstrator stages, preliminary surveys blank tests 1, 2 and 3.
- Fig. 4. Bentonite pellets and powder particle size distribution (Garitte et al. 2015).
- **Fig. 5.** Cross section of *Engineered Barrier System setup*.
- Fig. 6. Scheme of electrodes setup used for blank test 2 and 3 developed by IRSN.
- Fig. 7. (a) ERT protocol used for blank test 2 with ARES II unit in-line array, (b) ERT protocol
- used for blank test 2 with TERRAMETER LS cross-borehole array AM-BN (c) ERT protocol
- used for blank test 2 with TERRAMETER LS cross-borehole array AB-MN, where A and B are
- current electrodes and M and N are potential electrodes.
- Fig. 8. Electrodes resistance contacts (a) blank test 2 and (b) blank test 3.
- 573 **Fig. 9.** Stacking errors (a) blank test 2 and (b) blank test 3.
- Fig. 10. Distribution of apparent resistivity before and after filtering out measurements
- associated with large geometric factors, black and grey bars respectively, for blank test 2.
- 576 **Fig. 11.** Blank test 1: 2D surface survey, Schlumberger array (GN1 and WT-1 indicated by black
- 577 rectangles and main shaft and electrodes boreholes of ERT demonstrator area indicated by
- 578 black dashed rectangles).
- 579 **Fig. 12.** Blank test 2: borehole survey.
- Fig. 13. Blank test 3: cross borehole survey (buffer shaft indicated by black rectangle).
- Fig. 14. Synthetic data analysis. (a) 3D model (Model 2.4 x 2.6 x 10m, shaft 0.6 x 0.6 x 9.05m)
- (b) Perspective and cross section (at same plane where electrode boreholes are) view of 3D
- Data inverted using 3D algorithm and the AM-BN protocol of Blank test 3 and (c) 3D Data
- inverted using 2D algorithm and the AM-BN protocol of Blank test 3 (where A and B are
- current electrodes and M and N potential electrodes).

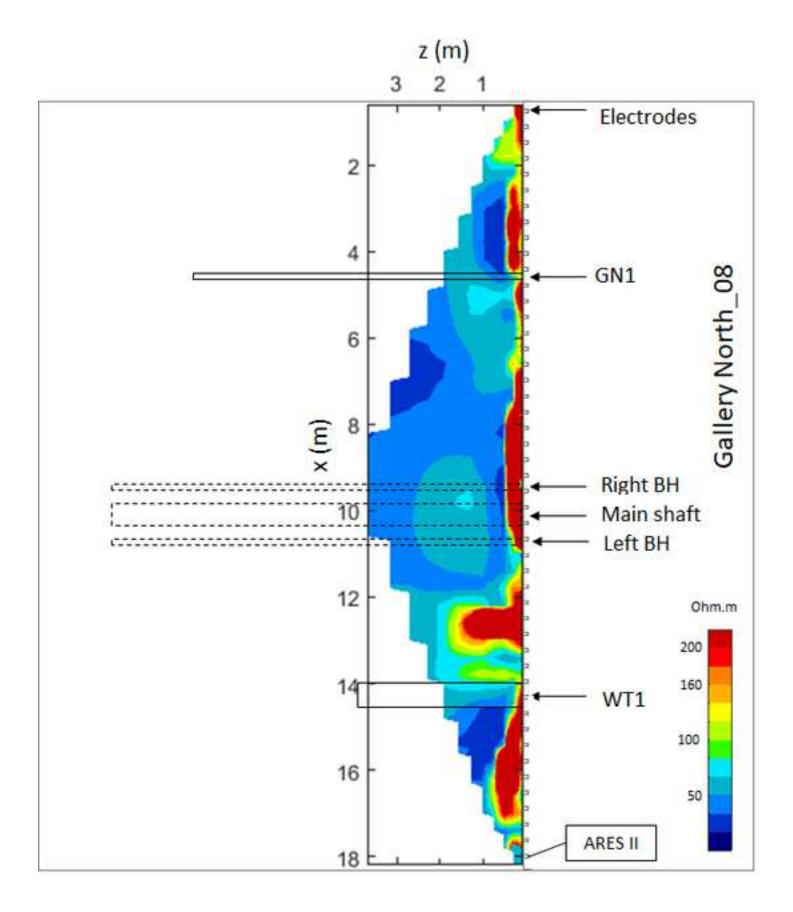
Table 1. Main characteristics of blank test 1.

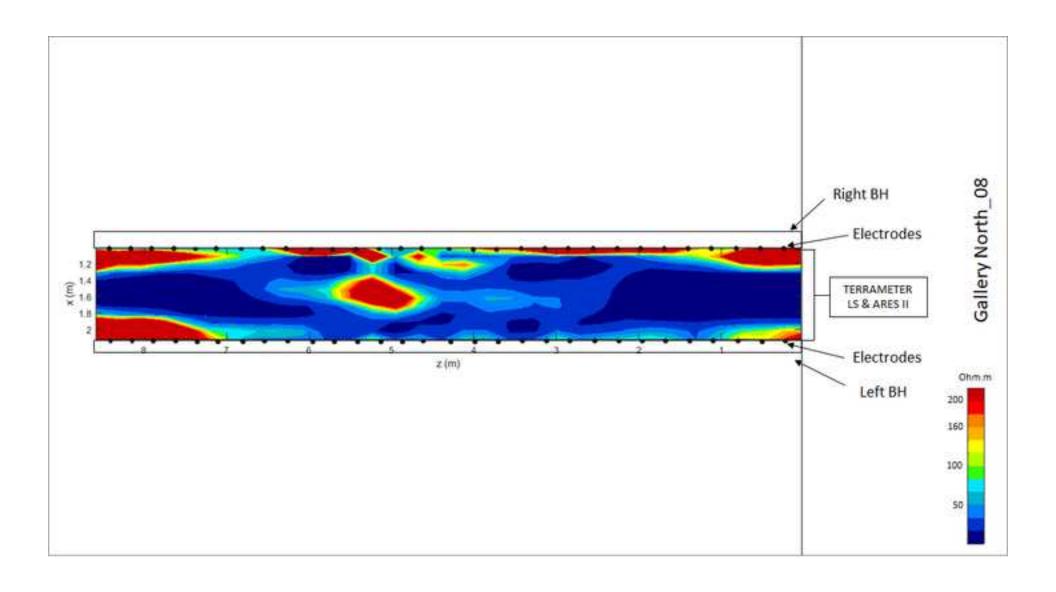
Electrodes spacing	0.4m		
Total number of	48		
electrodes			
Total length	18.8m		
Position of electrodes in	1.4m from the gallery floor		
z-axis			
First electrode (El 0) in x-	On the right: standing on Gallery North_08 and facing the ERT		
axis	demonstrator location		
Measurement type	2D surface		
Unit used	ARES II		
Array used	Schlumberger		
Electrodes used	Conventional metal sticks (surface)		

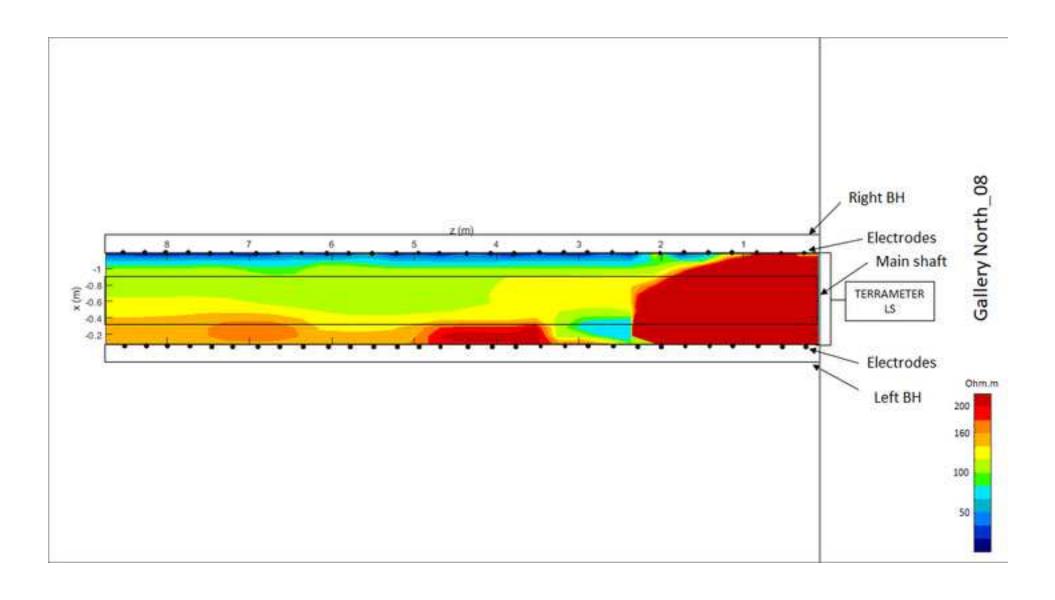
Table 2. Summary of number of data collected, stacking errors recorded, percentage of data removed in all blank tests and RMS errors obtained from inversions performed on original and filtered data sets.

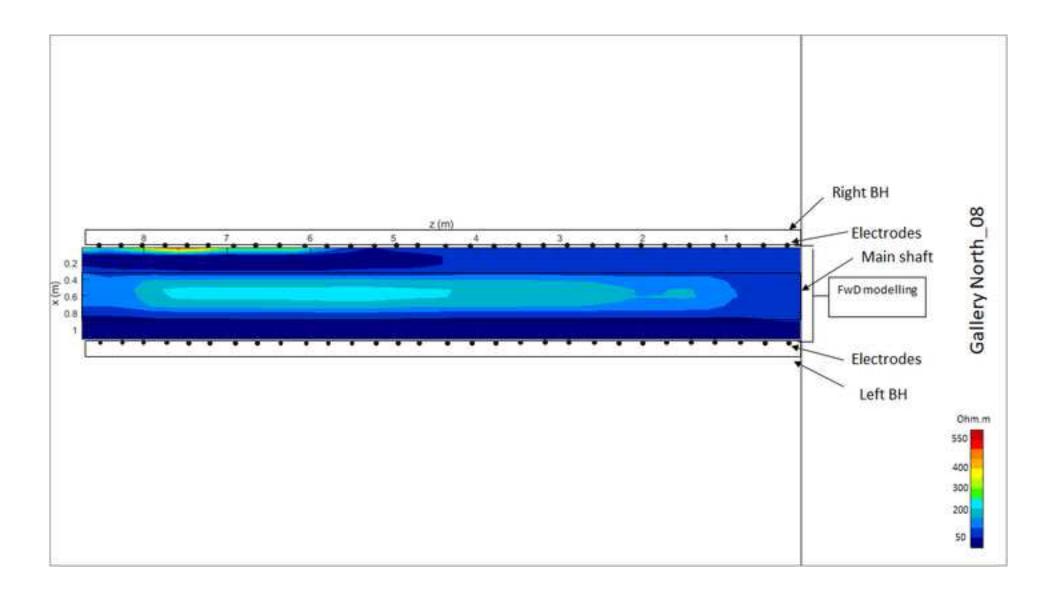
Surveys	Total No. of data	Mean stacking error (%)	Data stacking error > 3% (%)	Data removed (%)	Original data RMS (%)	Filtered data RMS (%)
Blank test 1	522	0.16	0.0	0.0	9.20	-
Blank test 2	1831	6.91	18.51	46.0	30.72	12.69
Blank test 3	1059	2.22	13.4	14.5	7.18	5.64

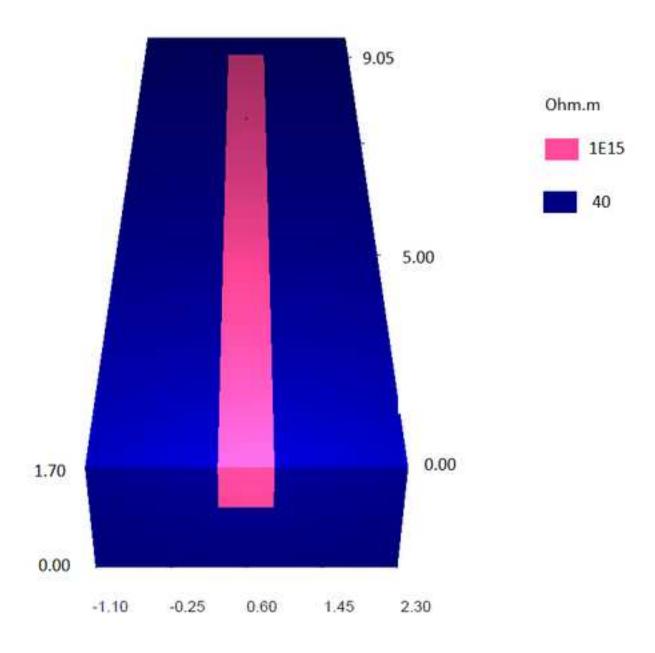


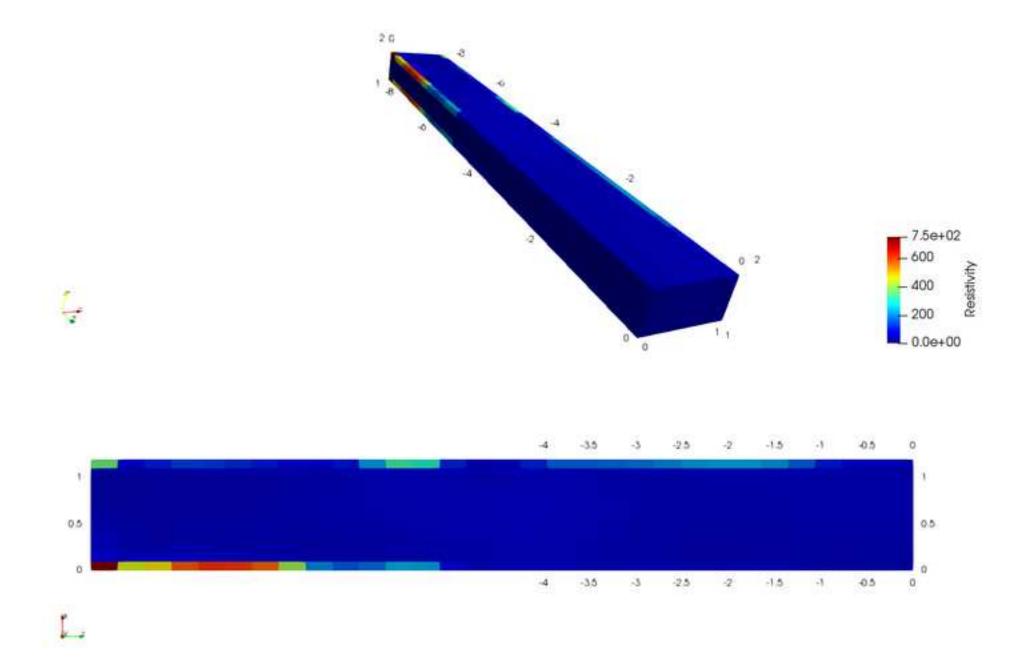


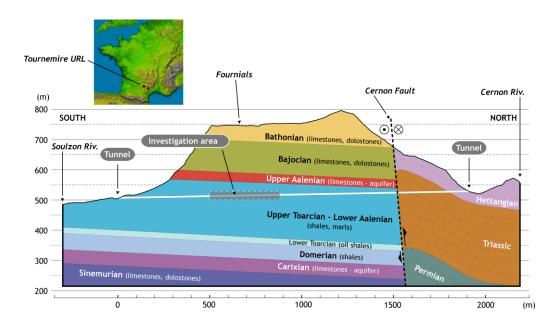


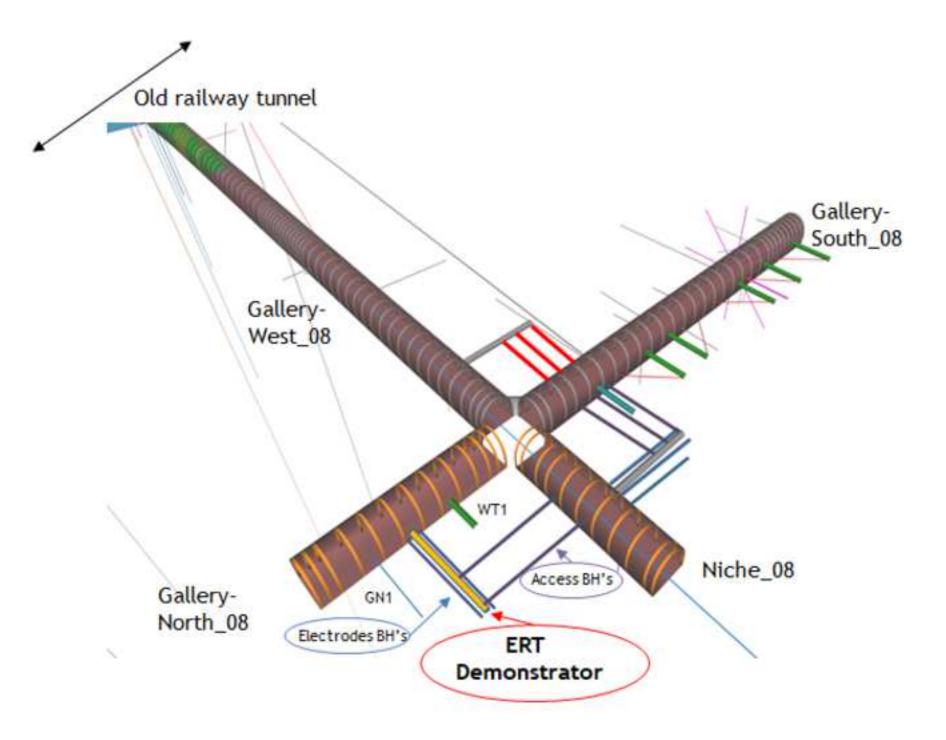




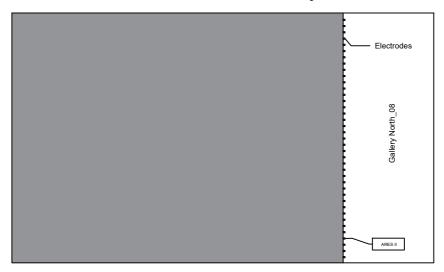




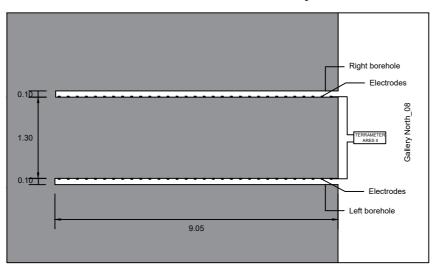




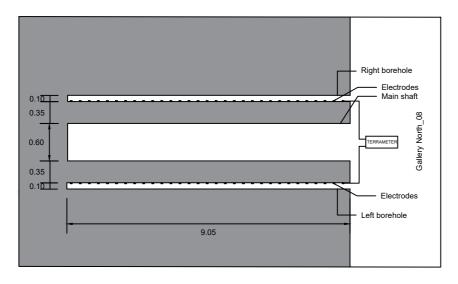
BLANK TEST 1 - January 2017

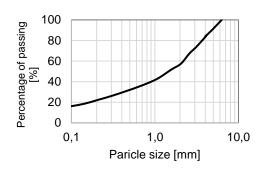


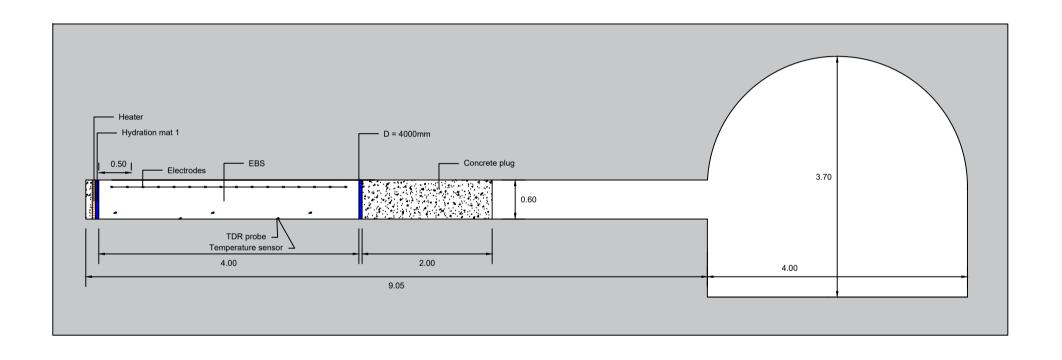
BLANK TEST 2 - January 2017

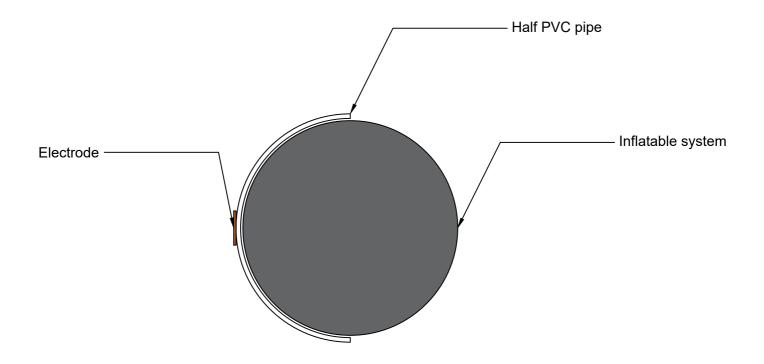


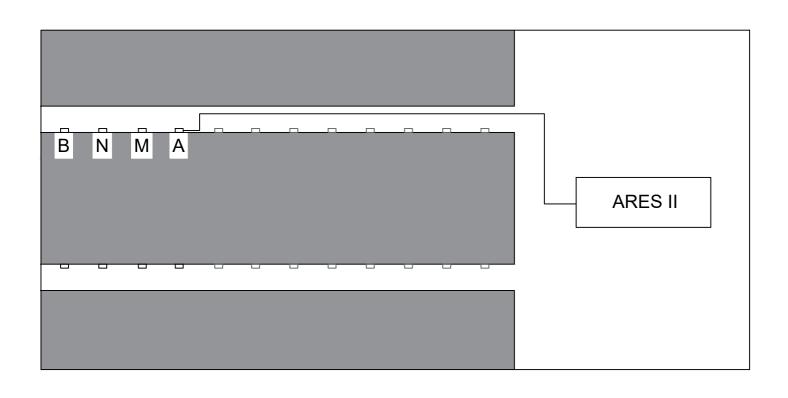
BLANK TEST 3 - November 2017

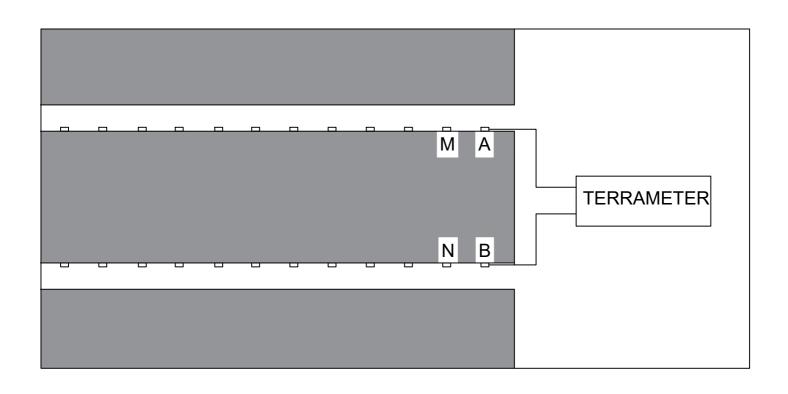


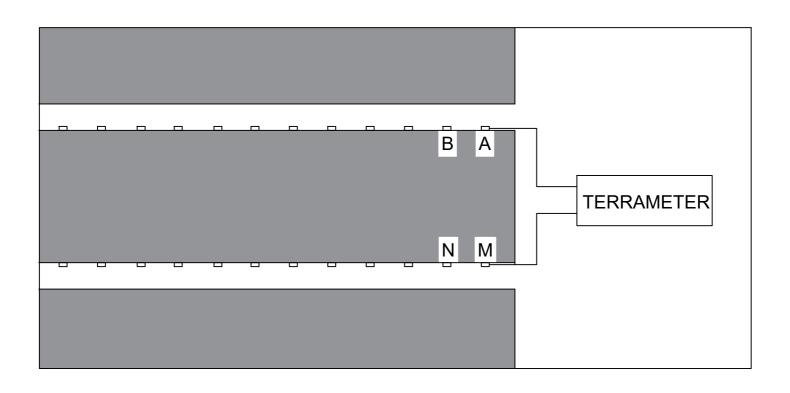


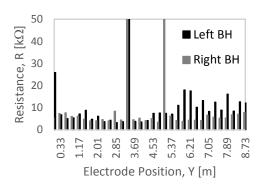


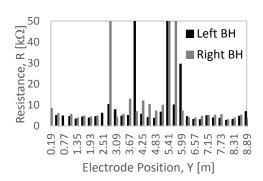


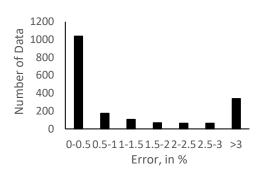


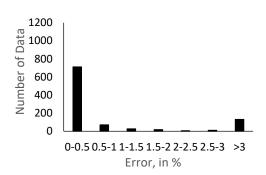


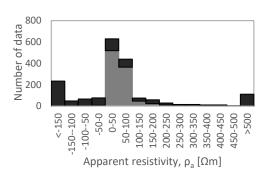












supplementary material (not datasets)

Click here to access/download supplementary material (not datasets)
Lopes_revised_tracked.docx