

Earthquake forecasting and its verification

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Abstract. No proven method is currently available for the reliable short time prediction of earthquakes (minutes to months). However, it is possible to make probabilistic hazard assessments for earthquake risk. In this paper we discuss a new approach to earthquake forecasting based on a pattern informatics (PI) method which quantifies temporal variations in seismicity. The output, which is based on an association of small earthquakes with future large earthquakes, is a map of areas in a seismogenic region (“hotspots”) where earthquakes are forecast to occur in a future 10-year time span. This approach has been successfully applied to California, to Japan, and on a worldwide basis. Because a sharp decision threshold is used, these forecasts are binary—an earthquake is forecast either to occur or to not occur. The standard approach to the evaluation of a binary forecast is the use of the relative (or receiver) operating characteristic (ROC) diagram, which is a more restrictive test and less subject to bias than maximum likelihood tests. To test our PI method, we made two types of retrospective forecasts for California. The first is the PI method and the second is a relative intensity (RI) forecast based on the hypothesis that future large earthquakes will occur where most smaller earthquakes have occurred in the recent past. While both retrospective forecasts are for the ten year period 1 January 2000 to 31 December 2009, we performed an interim analysis 5 years into the forecast. The PI method out performs the RI method under most circumstances.

1 Introduction

Earthquakes are the most feared of natural hazards because they occur without warning. Hurricanes can be tracked, floods develop gradually, and volcanic eruptions are preceded by a variety of precursory phenomena. Earthquakes, however, generally occur without any warning. There have

been a wide variety of approaches applied to the forecasting of earthquakes (Mogi, 1985; Turcotte, 1991; Lomnitz, 1994; Keilis-Borok, 2002; Scholz, 2002; Kanamori, 2003). These approaches can be divided into two general classes; the first is based on empirical observations of precursory changes. Examples include precursory seismic activity, precursory ground motions, and many others. The second approach is based on statistical patterns of seismicity. Neither approach has been able to provide reliable short-term forecasts (days to months) on a consistent basis.

Although short-term predictions are not available, long-term seismic-hazard assessments can be made. A large fraction of all earthquakes occur in the vicinity of plate boundaries, although some do occur in plate interiors. It is also possible to assess the long-term probability of having an earthquake of a specified magnitude in a specified region. These assessments are primarily based on the hypothesis that future earthquakes will occur in regions where past earthquakes have occurred (Frankel, 1995; Kossobokov et al., 2000). Specifically, the rate of occurrence of small earthquakes in a region can be analyzed to assess the probability of occurrence of much larger earthquakes.

The principal focus of this paper is a new approach to earthquake forecasting (Rundle et al., 2002; Tiampo et al., 2002b,a; Rundle et al., 2003; Appendix C). Our method does not predict the exact times and locations of earthquakes, but it does forecast the regions (hotspots) where earthquakes are most likely to occur in the relatively near future (typically ten years). The objective is to reduce the areas of earthquake risk relative to those given by long-term hazard assessments. Our approach is based on pattern informatics (PI), a technique that quantifies temporal variations in seismicity patterns. The result is a map of areas in a seismogenic region (hotspots) where earthquakes are likely to occur during a specified period in the future. A forecast for California was published by our group in 2002 (Rundle et al., 2002). Subsequently, sixteen of the eighteen California earthquakes with magnitudes $M \geq 5$ occurred in or immediately adjacent to the resulting hotspots. A forecast for Japan, presented in

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Tokyo in early October 2004, successfully forecast the location of the $M=6.8$ Niigata earthquake that occurred on 23 October 2004. A global forecast, presented at the early December 2004 meeting of the American Geophysical Union, successfully forecast the locations of the 23 December 2004, $M=8.1$ Macquarie Island earthquake, and the 26 December 2004 $M=9.0$ Sumatra earthquake. Before presenting further details of these studies we will give a brief overview of the current state of earthquake prediction and forecasting.

2 Empirical approaches

Empirical approaches to earthquake prediction rely on local observations of some type of precursory phenomena. In the near vicinity of the earthquake to be predicted, it has been suggested that one or more of the following phenomena may indicate a future earthquake (Mogi, 1985; Turcotte, 1991; Lomnitz, 1994; Keilis-Borok, 2002; Scholz, 2002; Kanamori, 2003): 1) precursory increase or decrease in seismicity in the vicinity of the origin of a future earthquake rupture, 2) precursory fault slip that leads to surface tilt and/or displacements, 3) electromagnetic signals, 4) chemical emissions, and 5) changes in animal behavior. Similarly, it has been suggested that precursory increases in seismic activity over large regions may indicate a future earthquake as well (Prozoroff, 1975; Dobrovolsky et al., 1979; Keilis-Borok et al., 1980; Press and Allen, 1995).

Examples of successful near-term predictions of future earthquakes have been rare. A notable exception was the prediction of the $M=7.3$ Haicheng earthquake in northeast China that occurred on 4 February 1975. This prediction led to the evacuation of the city which undoubtedly saved many lives. The Chinese reported that the successful prediction was based on foreshocks, groundwater anomalies, and animal behavior. Unfortunately, a similar prediction was not made prior to the magnitude $M=7.8$ Tangshan earthquake that occurred on 28 July 1976 (Utsu, 2003). Official reports placed the death toll in this earthquake at 242 000, although unofficial reports placed it as high as 655 000.

In order to thoroughly test for the occurrence of direct precursors the United States Geological Survey (USGS) initiated the Parkfield (California) Earthquake Prediction Experiment in 1985 (Bakun and Lindh, 1985; Kanamori, 2003). Earthquakes on this section of the San Andreas had occurred in 1857, 1881, 1901, 1922, 1934, and 1966. It was expected that the next earthquake in this sequence would occur by the early 1990's, and an extensive range of instrumentation was installed. The next earthquake in the sequence finally occurred on 28 September 2004. No precursory phenomena were observed that were significantly above the background noise level. Although the use of empirical precursors cannot be ruled out, the future of those approaches does not appear to be promising at this time.

3 Statistical and statistical physics approaches

A variety of studies have utilized variations in seismicity over relatively large distances to forecast future earthquakes. The distances are large relative to the rupture dimension of the subsequent earthquake. These approaches are based on the concept that the earth's crust is an activated thermodynamic system (Rundle et al., 2003). Among the evidence for this behavior is the continuous level of background seismicity in all seismographic areas. About a million magnitude two earthquakes occur each year on our planet. In southern California about a thousand magnitude two earthquakes occur each year. Except for the aftershocks of large earthquakes, such as the 1992 $M=7.3$ Landers earthquake, this seismic activity is essentially constant over time. If the level of background seismicity varied systematically with the occurrence of large earthquakes, earthquake forecasting would be relatively easy. This, however, is not the case.

There is increasing evidence that there are systematic precursory variations in some aspects of regional seismicity. For example, it has been observed that there is a systematic variation in the number of magnitude $M=3$ and larger earthquakes prior to at least some magnitude $M=5$ and larger earthquakes, and a systematic variation in the number of magnitude $M=5$ and larger earthquakes prior to some magnitude $M=7$ and larger earthquakes. The spatial regions associated with this phenomena tend to be relatively large, suggesting that an earthquake may resemble a phase change with an increase in the "correlation length" prior to an earthquake (Bowman et al., 1998; Jaumé and Sykes, 1999). There have also been reports of anomalous quiescence in the source region prior to a large earthquake, a pattern that is often called a "Mogi Donut" (Mogi, 1985; Kanamori, 2003; Wyss and Habermann, 1988; Wyss, 1997).

Many authors have noted the occurrence of a relatively large number of intermediate-sized earthquakes (foreshocks) prior to a great earthquake. A specific example was the sequence of earthquakes that preceded the 1906 San Francisco earthquake (Sykes and Jaumé, 1990). This seismic activation has been quantified as a power law increase in seismicity prior to earthquakes (Bowman et al., 1998; Jaumé and Sykes, 1999; Bufe and Varnes, 1993; Bufe et al., 1994; Brehm and Braile, 1998, 1999; Main, 1999; Robinson, 2000; Bowman and King, 2001; Yang et al., 2001; King and Bowman, 2003; Bowman and Sammis, 2004; Sammis et al., 2004). Unfortunately the success of these studies has depended on knowing the location of the subsequent earthquake.

A series of statistical algorithms to make intermediate term earthquake predictions have been developed by a Russian group under the direction of V. I. Keilis-Borok using pattern recognition techniques (Keilis-Borok, 1990, 1996). Seismicity in various circular regions was analyzed. Based primarily on seismic activation, earthquake alarms were issued for one or more regions, with the alarms generally lasting for five years. Alarms have been issued regularly since the mid 1980's and scored two notable successes: the prediction of the 1988 Armenian earthquake and the 1989 Loma Prieta

earthquake. While a reasonably high success rate has been achieved, there have been some notable misses including the recent $M=9.0$ Sumatra and $M=8.1$ Macquerie Island earthquakes.

More recently, this group has used chains of premonitory earthquakes as the basis for issuing alarms (Shebalin et al., 2004; Keilis-Borok et al., 2004). This method successfully predicted the $M=6.5$, 22 December 2003 San Simeon (California) earthquake and the $M=8.1$, 25 September 2003 Tokachi-oki, (Japan) earthquake with lead times of six and seven months respectively. However, an alarm issued for southern California, valid during the spring and summer of 2004, was a false alarm.

4 Chaos and forecasting

Earthquakes are caused by displacements on preexisting faults. Most earthquakes occur at or near the boundaries between the near-rigid plates of plate tectonics. Earthquakes in California are associated with the relative motion between the Pacific plate and the North American plate. Much of this motion is taken up by displacements on the San Andreas fault, but deformation and earthquakes extend from the Rocky Mountains on the east into the Pacific Ocean adjacent to California on the west. Clearly this deformation and the associated earthquakes are extremely complex.

Slider-block models are considered to be simple analogs to seismicity. A pair of interacting slider-blocks have been shown to exhibit classical chaotic behavior (Turcotte, 1997). This low-order behavior is taken to be evidence for the chaotic behavior of seismicity in the same way that the chaotic behavior of the low-order Lorentz equations is taken as evidence for the chaotic behavior of weather and climate. Some authors (Geller et al., 1997; Geller, 1997) have argued that this chaotic behavior precludes the prediction of earthquakes. However, weather is also chaotic, but forecasts can be made. Weather forecasts are probabilistic in the sense that weather cannot be predicted exactly. One such example is the track of a hurricane. Probabilistic forecasts of hurricane tracks are routinely made; sometimes they are extremely accurate while at other times they are not. Another example of weather forecasting is the forecast of El Niño events. Forecasting techniques based on pattern recognition and principle components of the sea surface temperature fluctuation time series have been developed that are quite successful in forecasting future El Niños, but again they are probabilistic in nature (Chen et al., 2004). It has also been argued (Sykes et al., 1999) that chaotic behavior does not preclude the probabilistic forecasting of future earthquakes. Over the past five years our group has developed (Rundle et al., 2002; Tiampo et al., 2002b,a; Rundle et al., 2003; Holliday et al., 2005a) a technique for forecasting the locations where earthquakes will occur based on pattern informatics (PI). This type of approach has close links to principle component analysis, which has been successfully used for the forecasting of El Niños.

5 The PI method

Seismic networks provide the times and locations of earthquakes over a wide range of scales. One of the most sensitive networks has been deployed over southern California and the resulting catalog is readily available. Our objective has been to analyze the historical seismicity for anomalous behavior that would provide information on the occurrence of future earthquakes. At this point we are not able to forecast the times of future earthquakes with precision. However, our approach does appear to select the regions where earthquakes are most likely to occur during a future time window. At the present time, this time window is typically taken to be ten years, although it appears that it is possible to utilize shorter time windows.

Our approach divides the seismogenic region to be studied into a grid of square boxes whose size is related to the magnitude of the earthquakes to be forecast. The rates of seismicity in each box are studied to quantify anomalous behavior. The basic idea is that any seismicity precursors represent changes, either a local increase or decrease of seismic activity, so our method identifies the locations in which these changes are most significant during a predefined change interval. The subsequent forecast interval is the decadal time window during which the forecast is valid. The box size is selected to be consistent with the size of the earthquakes being forecasted (Bowman et al., 1998), and the minimum earthquake magnitude considered is the lower limit of sensitivity and completeness of the network in the region under consideration.

While a detailed explanation of the PI method that we have used for earthquake forecasting is included in Appendix A, a compact utilization is given as follows:

1. The region of interest is divided into N_B square boxes with linear dimension Δx . Boxes are identified by a subscript i and are centered at x_i . For each box, there is a time series $N_i(t)$, which is the instantaneous number of earthquakes per unit time at time t larger than the lower cut-off magnitude M_c . The time series in box i is defined between a base time t_b and the present time t .
2. All earthquakes in the region of interest with magnitudes greater than a lower cutoff magnitude M_c are included. The lower cutoff magnitude M_c is specified in order to ensure completeness of the data through time, from an initial time t_0 to a final time t_2 .
3. Three time intervals are considered:
 - (a) A reference time interval from t_b to t_1 .
 - (b) A second time interval from t_b to t_2 , $t_2 > t_1$. The change interval over which seismic activity changes are determined is then $t_2 - t_1$. The time t_b is chosen to lie between t_0 and t_1 . Typically we take $t_0=1932$, $t_1=1990$, and $t_2=2000$. The objective is to quantify anomalous seismic activity in the change interval t_1 to t_2 relative to the reference interval t_b to t_1 .

(c) The forecast time interval t_2 to t_3 , for which the forecast is valid. We take the change and forecast intervals to have the same length. For the above example, $t_3=2010$.

4. The seismic intensity in box i , $I_i(t_b, t)$, between two times $t_b < t$, can then be defined as the average number of earthquakes with magnitudes greater than M_c that occur in the box per unit time during the specified time interval t_b to t . Therefore, using discrete notation, we can write:

$$I_i(t_b, t) = \frac{1}{t - t_b} \sum_{t'=t_b}^t N_i(t'), \quad (1)$$

where the sum is performed over increments of the time series, say days.

5. In order to compare the intensities from different time intervals, we require that they have the same statistical properties. We therefore normalize the seismic intensities by subtracting the mean seismic activity of all boxes and dividing by the standard deviation of the seismic activity in all boxes. The statistically normalized seismic intensity of box i during the time interval t_b to t is then defined by

$$\hat{I}_i(t_b, t) = \frac{I_i(t_b, t) - \langle I(t_b, t) \rangle}{\sigma(t_b, t)}, \quad (2)$$

where $\langle I(t_b, t) \rangle$ is the mean intensity averaged over all the boxes and $\sigma(t_b, t)$ is the standard deviation of intensity over all the boxes.

6. Our measure of anomalous seismicity in box i is the difference between the two normalized seismic intensities:

$$\Delta I_i(t_b, t_1, t_2) = \hat{I}_i(t_b, t_2) - \hat{I}_i(t_b, t_1). \quad (3)$$

This measure is motivated by the assumption of pure phase dynamics (Rundle et al., 2000a,b) that important changes in seismicity will be given by the change in the anomalous seismicity over time.

7. To reduce the relative importance of random fluctuations (noise) in seismic activity, we compute the average change in intensity, $\overline{\Delta I_i(t_0, t_1, t_2)}$ over all possible pairs of normalized intensity maps having the same change interval:

$$\overline{\Delta I_i(t_0, t_1, t_2)} = \frac{1}{t_1 - t_0} \sum_{t_b=t_0}^{t_1} \Delta I_i(t_b, t_1, t_2), \quad (4)$$

where the sum is performed over increments of the time series, which here are days.

8. We hypothesize that the probability of a future earthquake in box i , $P_i(t_0, t_1, t_2,)$, is proportional to the square of the average intensity change:

$$P_i(t_0, t_1, t_2,) \propto \overline{\Delta I_i(t_b, t_1, t_2)}^2. \quad (5)$$

The constant of proportionality can be determined by requiring unit probability but is not important to the analysis.

9. To identify anomalous regions, we wish to compute the change in the probability $P_i(t_0, t_1, t_2,)$ relative to the background so that we subtract the mean probability over all boxes. We denote this change in the probability by

$$\Delta P_i(t_0, t_1, t_2) = P_i(t_0, t_1, t_2) - \langle P_i(t_0, t_1, t_2) \rangle, \quad (6)$$

where $\langle P_i(t_0, t_1, t_2) \rangle$ is the background probability for a large earthquake.

Hotspots are defined to be the regions where $\Delta P_i(t_0, t_1, t_2)$ is positive. In these regions, $P_i(t_0, t_1, t_2)$ is larger than the average value for all boxes (the background level). Note that since the intensities are squared in defining probabilities the hotspots may be due to either increases of seismic activity during the change time interval (activation) or due to decreases (quiescence). We hypothesize that earthquakes with magnitudes larger than M_c+2 will occur preferentially in hotspots during the forecast time interval t_2 to t_3 .

6 Applications of the PI method

The PI method was first applied to seismicity in southern California and adjacent regions (32° to 37° N lat, 238° to 245° E long). This region was divided into a grid of 3500 boxes with $\Delta x=0.1^\circ$ (11 km). Consistent with the sensitivity of the southern California seismic network, the lower magnitude cutoff was taken to be $M=3$. The initial time was $t_0=1932$, the change interval was from $t_1=1990$ to $t_2=2000$, and the forecast interval was from $t_2=2000$ to $t_3=2010$. The initial studies for California were published in 2002 (Rundle et al., 2002), the results are reproduced in Fig. 1. The colored regions are the hotspots defined to be the boxes where ΔP is positive. This forecast of where earthquakes would likely occur was considered to be valid for the forecast interval from 2000 to 2010 and would be applicable for earthquakes with $M=5$ and larger. Since 1 January 2000, eighteen significant earthquakes have occurred in the test region. We consider a significant earthquake to be an event that was initially assigned a magnitude 5 or larger. These are also shown in Fig. 1, and information on these earthquakes is given in Table 1. We consider the forecast to be successful if the epicenter of the earthquake lies within a hotspot box or in one of the eight adjoining boxes (Moore, 1962). Sixteen of the eighteen earthquakes were successfully forecast.

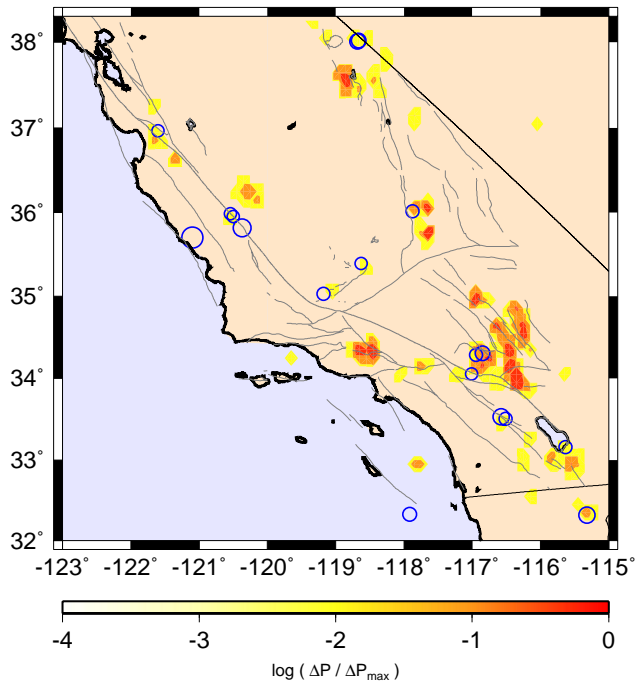


Fig. 1. Application of the PI method to southern California. Colored areas are the forecast hotspots for the occurrence of $M \geq 5$ earthquakes during the period 2000–2010 derived using the PI method. The color scale gives values of $\log_{10}(P/P_{max})$. Also shown are the locations of the eighteen earthquakes with $M \geq 5$ that have occurred in the region since 1 January 2000. Sixteen of the eighteen earthquakes were successfully forecast. More details of the earthquakes are given in Table 1.

The second area to which the PI method was applied was Japan. The forecast hotspots for the Tokyo region (33° to 38° N lat, 136° to 142° W long) are given in Fig. 2. The initial time was $t_0=1965$, and the change and forecast intervals were the same as those used for California. Between 1 January 2000 and 14 October 2004, 99 earthquakes occurred and 91 earthquakes were successfully forecast. This forecast was presented at the International Conference on Geodynamics, 14–16 October 2004, Tokyo by one of the authors (JBR). Subsequently the Niigata earthquake ($M=6.8$) occurred on 23 October 2004. This earthquake and its subsequent $M \geq 5$ aftershocks were successfully forecast.

The PI method has also been applied on a worldwide basis. In this case $1^\circ \times 1^\circ$ boxes were considered, $\Delta x=1^\circ$ (110 km). Consistent with the sensitivity of the global seismic network the lower magnitude cutoff was taken to be $M_c = 5$. The initial time was $t_0=1965$; the change and forecast intervals were the same as above. The resulting map of hotspots was presented by two of the authors (DLT and JRH) at the Fall Meeting of the American Geophysical Union on 14 December 2004 (abstracts: AGUF2004NG24B-01 and AGUF2004NG54A-08). This map is given in Fig. 3. This forecast of where earthquakes would occur was considered to be valid for the period 2000 to 2010 and would be applicable for earthquakes with magnitudes greater than 7.0.

Table 1. Earthquakes with $M \geq 5$ that occurred in the California test region since 1 January 2000. Sixteen of these eighteen earthquakes were successfully forecast. The two missed events are marked with an asterisk.

	Event	Magnitude	Local Time
1	Big Bear I	$M=5.1$	10 Feb. 2001
2	Coso	$M=5.1$	17 July 2001
3	Anza I	$M=5.1$	31 Oct. 2001
4	Baja	$M=5.7$	22 Feb. 2002
5	Gilroy	$M=5.0$	14 May 2002
6	Big Bear II	$M=5.4$	22 Feb. 2003
7	San Simeon*	$M=6.5$	22 Dec. 2003
8	San Clemente Island*	$M=5.2$	15 June 2004
9	Bodie I	$M=5.5$	18 Sep. 2004
10	Bodie II	$M=5.4$	18 Sep. 2004
11	Parkfield I	$M=6.0$	28 Sep. 2004
12	Parkfield II	$M=5.2$	29 Sep. 2004
13	Arvin	$M=5.0$	29 Sep. 2004
14	Parkfield III	$M=5.0$	30 Sep. 2004
15	Wheeler Ridge	$M=5.2$	16 April 2005
16	Anza II	$M=5.2$	12 June 2005
17	Yucaipa	$M=4.9$	16 June 2005
18	Obsidian Butte	$M=5.1$	2 Sep. 2005

Between 1 January 2000 and 14 December 2004 there were sixty eight $M \geq 7$ earthquakes worldwide; fifty seven of these earthquakes occurred within a hotspot or adjoining boxes. Subsequent to the meeting presentation, the $M=8.1$ Macquarie Island earthquake occurred on 23 December 2004 and the $M=9.0$ Sumatra earthquake occurred on 26 December 2004. The epicenters of both earthquakes were successfully forecast.

7 Forecast verification

Previous tests of earthquake forecasts have emphasized the likelihood test (Kagan and Jackson, 2000; Rundle et al., 2002; Tiampo et al., 2002b; Holliday et al., 2005a). These tests have the significant disadvantage that they are overly sensitive to the least probable events. For example, consider two forecasts. The first perfectly forecasts 99 out of 100 events but assigns zero probability to the last event. The second assigns zero probability to all 100 events. Under a log-likelihood test, both forecasts will have the same skill score of $-\infty$. Furthermore, a naive forecast that assigns uniform probability to all possible sites will always score higher than a forecast that misses only a single event but is otherwise superior. For this reason, likelihood tests are more subject to unconscious bias. Other methods of evaluating earthquake forecasts are suggested by Harte and Vere-Jones (2005) and Holliday et al. (2005b).

An extensive review on forecast verification in the atmospheric sciences has been given by Jolliffe and Stephenson (2003). The wide variety of approaches that they consider

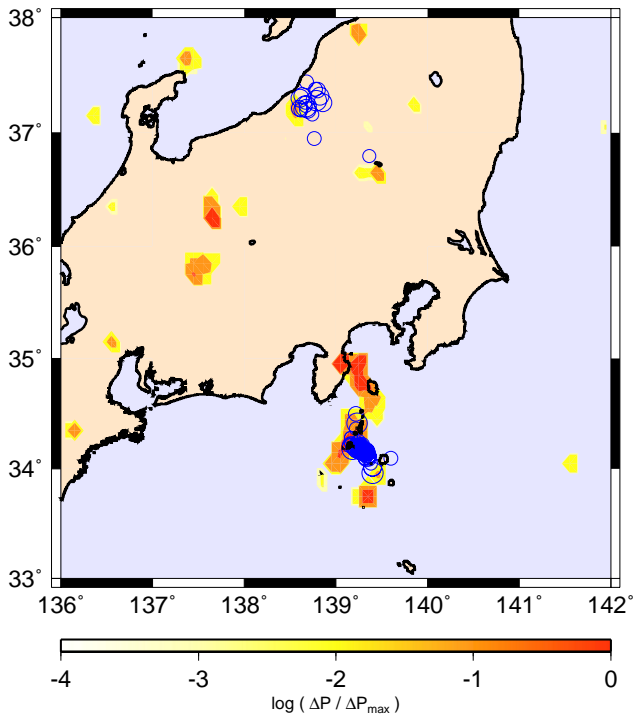


Fig. 2. Application of the PI method to central Japan. Colored areas are the forecast hotspots for the occurrence of $M \geq 5$ earthquakes during the period 2000-2010 derived using the PI method. The color scale gives values of the $\log_{10}(P/P_{\max})$. Also shown are the locations of the 99 earthquakes with $M \geq 5$ that have occurred in the region since 1 January 2000.

are directly applicable to earthquake forecasts as well. The earthquake forecasts considered in this paper can be viewed as binary forecasts by considering the events (earthquakes) as being forecast either to occur or not to occur in a given box. We consider that there are four possible outcomes for each box, thus two ways to classify each red, hotspot, box, and two ways to classify each white, non-hotspot, box:

1. An event occurs in a hotspot box or within the Moore neighborhood of the box (the Moore neighborhood is comprised of the eight boxes surrounding the forecast box). This is a success.
2. No event occurs in a white non-hotspot box. This is also a success.
3. No event occurs in a hotspot box or within the Moore neighborhood of the hotspot box. This is a false alarm.
4. An event occurs in a white, non-hotspot box that is not within the Moore neighborhood of a hotspot box. This is a failure to forecast.

We note that these rules tend to give credit, as successful forecasts, for events that occur very near hotspot boxes. We have adopted these rules in part because the grid of boxes is positioned arbitrarily on the seismically active region, thus

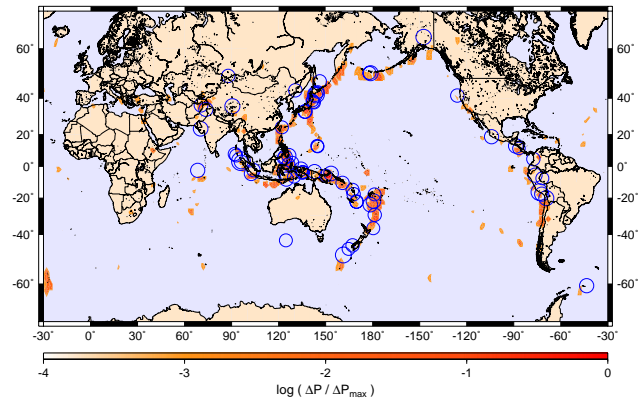


Fig. 3. World-wide application of the PI method. Colored areas are the forecast hotspots for the occurrence of $M \geq 7$ earthquakes during the period 2000-2010 derived using the PI method. The color scale gives values of the $\log_{10}(P/P_{\max})$. Also shown are the locations of the sixty eight earthquakes with $M \geq 7$ that have occurred in the region since 1 January 2000.

we allow a margin of error of ± 1 box dimension. In addition, the events we are forecasting are large enough so that their source dimension approaches, and can even exceed, the box dimension meaning that an event might have its epicenter outside a hotspot box, but the rupture might then propagate into the box. Other similar rules are possible but we have found that all such rules basically lead to similar results.

The standard approach to the evaluation of a binary forecast is the use of a relative operating characteristic (ROC) diagram (Swets, 1973; Mason, 2003). Standard ROC diagrams consider the fraction of failures-to-predict and the fraction of false alarms. This method evaluates the performance of the forecast method relative to random chance by constructing a plot of the fraction of failures to predict against the fraction of false alarms for an ensemble of forecasts. Molchan (1997) has used a modification of this method to evaluate the success of intermediate term earthquake forecasts.

The binary approach has a long history, over 100 years, in the verification of tornado forecasts (Mason, 2003). These forecasts take the form of a tornado forecast for a specific location and time interval, each forecast having a binary set of possible outcomes. For example, during a given time window of several hours duration, a forecast is issued in which a list of counties is given with a statement that one or more tornadoes will or will not occur. A 2×2 contingency table is then constructed, the top row contains the counties in which tornadoes are forecast to occur and the bottom row contains counties in which tornadoes are forecast to not occur. Similarly, the left column represents counties in which tornadoes were actually observed, and the right column represents counties in which no tornadoes were observed.

With respect to earthquakes, our forecasts take exactly this form. A time window is proposed during which the forecast of large earthquakes having a magnitude above some minimum threshold is considered valid. An example might be a forecast of earthquakes larger than $M=5$ during a period

of five or ten years duration. A map of the seismically active region is then completely covered (“tiled”) with boxes of two types: boxes in which the epicenters of at least one large earthquake are forecast to occur and boxes in which large earthquakes are forecast to not occur. In other types of forecasts, large earthquakes are given some continuous probability of occurrence from 0% to 100% in each box (Kagan and Jackson, 2000). These forecasts can be converted to the binary type by the application of a decision threshold. Boxes having a probability below the threshold are assigned a forecast rating of *non-occurrence* during the time window, while boxes having a probability above the threshold are assigned a forecast rating of *occurrence*. A high threshold value may lead to many *failures to predict* (events that occur where no event is forecast), but few *false alarms* (an event is forecast at a location but no event occurs). The level at which the decision threshold is set is then a matter of public policy specified by emergency planners, representing a balance between the prevalence of failures to predict and false alarms.

8 Binary earthquake forecast verification

To illustrate this approach to earthquake forecast verification, we have constructed two types of retrospective binary forecasts for California. The first type of forecast utilizes the PI method described above. We apply the method to southern California and adjacent regions (32° to 38.3° N lat, 238° to 245° E long) using a grid of boxes with $\Delta x=0.1^\circ$ and a lower magnitude cutoff $M_c=3.0$. For this retrospective forecast we take the initial time $t_0=1932$, the change interval $t_1=1989$ to $t_2=2000$, and the forecast interval $t_2=2000$ to $t_3=2010$ (Rundle et al., 2002; Tiampo et al., 2002b). In the analysis given above we considered regions with ΔP positive to be hotspots. The PI forecast under the above conditions with $\Delta P>0$ is given in Fig. 4b. Hotspots include 127 of the 5040 boxes considered. This forecast corresponds to that given in Fig. 1. The threshold for hotspot activation can be varied by changing the threshold value for ΔP . A forecast using a higher threshold value is given in Fig. 4a. Hotspots here include only 29 of the 5040 boxes considered.

An alternative approach to earthquake forecasting is to use the rate of occurrence of earthquakes in the past. We refer to this type of forecast as a *relative intensity* (RI) forecast. In such a forecast, the study region is tiled with boxes of size $0.1^\circ \times 0.1^\circ$. The number of earthquakes with magnitude $M \geq 3.0$ in each box down to a depth of 20 km is determined over the time period from $t_0=1932$ to $t_2=2000$. The RI score for each box is then computed as the total number of earthquakes in the box in the time period divided by the value for the box having the largest value. A threshold value in the interval $[0, 1]$ is then selected. Large earthquakes having $M \geq 5$ are then considered possible only in boxes having an RI value larger than the threshold. The remaining boxes with RI scores smaller than the threshold represent sites at which large earthquakes are forecast to not occur. The physical justification for this type of forecast is that large earthquakes

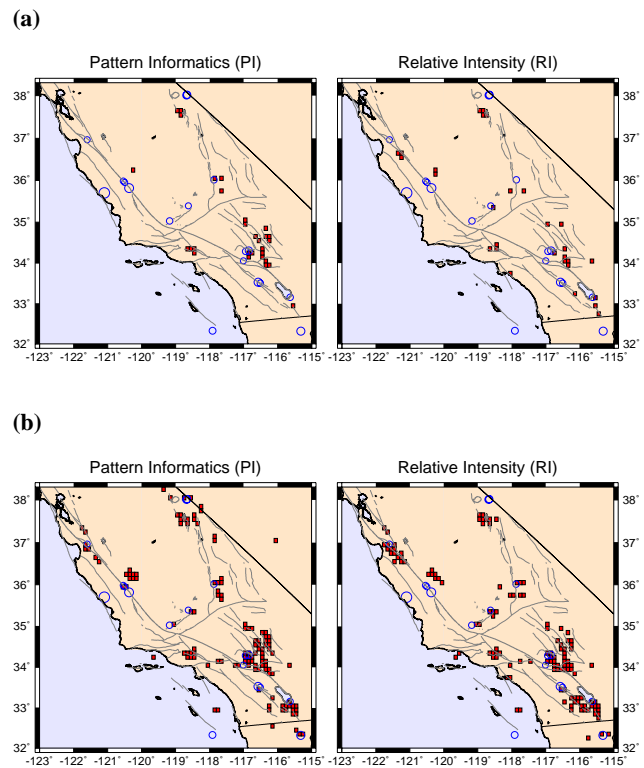


Fig. 4. Retrospective application of the PI and RI methods for southern California as a function of false alarm rate. Red boxes are the forecast hotspots for the occurrence of $M \geq 5$ earthquakes during the period 2000 to 2005. Also shown are the locations of the $M \geq 5$ earthquakes that occurred in this region during the forecast period. In (a), a threshold value was chosen such that $F \approx 0.005$. In (b), a threshold value was chosen such that $F \approx 0.021$.

are considered most likely to occur at sites of high seismic activity.

In order to make a direct comparison of the RI forecast with the PI forecast, we select the threshold for the RI forecast to give the same box coverage given for the PI forecast in Figs. 4a and 4b, i.e. 29 boxes and 127 boxes respectively. Included in all figures are the earthquakes with $M \geq 5$ that occurred between 2000 and 2005 in the region under consideration.

9 Contingency tables and ROC diagrams

The first step in our generation of ROC diagrams is the construction of the 2×2 contingency table for the PI and RI forecast maps given in Fig. 4. The hotspot boxes in each map represent the forecast locations. A hotspot box upon which *at least one* large future earthquake during the forecast period occurs is counted as a *successful forecast*. A hotspot box upon which *no* large future earthquake occurs during the forecast period is counted as an *unsuccessful forecast*, or alternately, a *false alarm*. A white box upon which *at least one* large future earthquake during the forecast period occurs is counted as a *failure to forecast*. A white box upon which *no*

Table 2. Contingency tables as a function of false alarm rate. In (a), a threshold value was chosen such that $F \approx 0.005$. In (b), a threshold value was chosen such that $F \approx 0.021$.

(a)

Pattern informatics (PI) forecast			
Forecast	Observed		Total
	Yes	No	
Yes	(a) 4	(b) 25	29
No	(c) 13	(d) 4998	5011
Total	17	5023	5040

Relative intensity (RI) forecast			
Forecast	Observed		Total
	Yes	No	
Yes	(a) 2	(b) 27	29
No	(c) 14	(d) 4997	5011
Total	16	5024	5040

(b)

Pattern informatics (PI) forecast			
Forecast	Observed		Total
	Yes	No	
Yes	(a) 23	(b) 104	127
No	(c) 9	(d) 4904	4913
Total	32	5008	5040

Relative intensity (RI) forecast			
Forecast	Observed		Total
	Yes	No	
Yes	(a) 20	(b) 107	127
No	(c) 10	(d) 4903	4913
Total	30	5010	5040

large future earthquake occurs during the forecast period is counted as a *successful forecast of non-occurrence*.

Verification of the PI and RI forecasts proceeds in exactly the same way as for tornado forecasts. For a given number of hotspot boxes, which is controlled by the value of the probability threshold in each map, the contingency table (see Table 2) is constructed for both the PI and RI maps. Values for the table elements a (Forecast=yes, Observed=yes), b (Forecast=yes, Observed=no), c (Forecast=no, Observed=yes), and d (Forecast=no, Observed=no) are obtained for each map. The fraction of colored boxes, also called the *probability of forecast of occurrence*, is $r=(a+b)/N$, where the total number of boxes is $N=a+b+c+d$. The *hit rate* is $H=a/(a+c)$ and is the fraction of large earthquakes that occur on a hotspot. The *false alarm rate* is $F=b/(b+d)$ and is the fraction of non-observed earthquakes that are incorrectly forecast.

To analyze the information in the PI and RI maps, the standard procedure is to consider all possible forecasts together. These are obtained by increasing F from 0 (corresponding to no hotspots on the map) to 1 (all active boxes on the map are identified as hotspots). The plot of H versus F is the relative operating characteristic (ROC) diagram. Varying the threshold value for both the PI and RI forecasts, we have obtained the values of H and F given in Fig. 5, blue for the PI forecasts and red for the RI forecasts. The results corresponding to the maps given in Fig. 4 and the contingency tables given in Table 2 are given by the filled symbols. The forecast with 29 hotspot boxes (Fig. 5a and Table 2a) has $F_{PI}=0.00498$, $H_{PI}=0.235$ and $F_{RI}=0.00537$, $H_{RI}=0.125$. The forecast with 127 hotspot boxes (Fig. 5b and Table 2b) has $F_{PI}=0.0207$, $H_{PI}=0.719$ and $F_{RI}=0.0213$, $H_{RI}=0.666$. Also shown in Fig. 5 is a gain curve (green) defined by the ratio of $H_{PI}(F)$ to $H_{RI}(F)$. Gain values greater than unity indicate better performance using the PI map than using the RI map. The horizontal dashed line corresponds to zero gain. From Fig. 5 it can be seen that the PI approach outperforms (is above) the RI under many circumstances and both outperform a random map, where $H=F$, by a large margin.

10 Discussion

The fundamental question is whether forecasts of the time and location of future earthquakes can be accurately made. It is accepted that long term hazard maps of the expected rate of occurrence of earthquakes are reasonably accurate. But is it possible to do better? Are there precursory phenomena that will allow earthquakes to be forecast?

It is actually quite surprising that immediate local precursory phenomena are not seen. Prior to a volcanic eruption, increases in regional seismicity and surface movements are generally observed. For a fault system, the stress gradually increases until it reaches the frictional strength of the fault and a rupture is initiated. It is certainly reasonable to hypothesize that the stress increase would cause increases in background seismicity and aseismic slip. In order to test this hypothesis the Parkfield Earthquake Prediction Experiment was initiated in 1985. The expected Parkfield earthquake occurred beneath the heavily instrumented region on 28 September 2004. No local precursory changes were observed (Lindh, 2005).

In the absence of local precursory signals, the next question is whether broader anomalies develop, and in particular whether there is anomalous seismic activity. It is this question that is addressed in this paper. Using a technique that has been successfully applied to the forecasting of El Niño we have developed a systematic pattern informatics (PI) approach to the identification of regions of anomalous seismic activity. Applications of this technique to California, Japan, and on a world-wide basis have successfully forecast the location of future earthquakes. It must be emphasized that this is not an earthquake prediction and does not state exactly

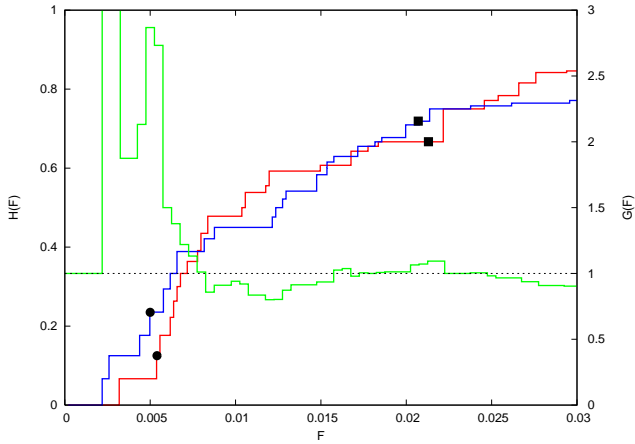


Fig. 5. Relative operating characteristic (ROC) diagram. Plot of hit rates, H , versus false alarm rates, F , for the PI forecast (blue) and RI forecast (red). Also shown is the gain ratio (green) defined as $H_{PI}(F)/H_{RI}(F)$. The filled symbols correspond to the threshold values used in Fig. 4 and Table 2, solid circles for 29 hotspot boxes and solid squares for 127 hotspot boxes. The horizontal dashed line corresponds to zero gain.

when and where the next earthquake will occur. It is a forecast of where future earthquakes are expected to preferentially occur during a relatively long time window of ten years. The objective is to reduce the possible future sites of earthquakes relative to a long term hazard assessment map.

Examination of the ROC diagrams indicates that the most important and useful of the suite of forecast maps are those with the least number of hotspot boxes, i.e. those with small values of the false alarm rate, F . A relatively high proportion of these hotspot boxes represent locations of future large earthquakes, however these maps also have a larger number of failures-to-forecast. Exactly which forecast map(s) to be used will be a decision for policy-makers, who will be called upon to balance the need for few false alarms against the desire for the least number of failures-to-forecast.

Finally, we remark that the methods used to produce the forecast maps described here can be extended and improved. In it's current state, the PI map is quite similar to the RI map. We have found modifications to the procedures described in Section 5 that allow the PI map to substantially outperform the RI map as indicated by the respective ROC diagrams. These methods are based on the approach of: 1) starting with the RI map and introducing improvements using the steps described for the PI method; and 2) introducing an additional averaging step. This new method is outlined in Appendix B, and a full analysis with results will be presented in a future publication.

Appendix A Explanation of the PI method

Here we summarize the current PI method as described by Rundle et al. (2003) and Tiampo et al. (2002b). The PI approach is a six step process that creates a time-dependent sys-

tem state vector in a real valued Hilbert space and uses the phase angle to predict future states (Rundle et al., 2003). The method is based on the idea that the future time evolution of seismicity can be described by pure phase dynamics (Mori and Kuramoto, 1998; Rundle et al., 2000a,b). Hence, a real-valued seismic phase function $\hat{I}_i(t_b, t)$ is constructed and allowed to rotate in its Hilbert space. Since seismicity in active regions is a noisy function (Kanamori, 1981), only temporal averages of seismic activity are utilized in the method. The geographic area of interest is partitioned into N square bins with an edge length δx determined by the nature of the physical system. For our analysis we chose $\delta x = 0.1^\circ \approx 11$ km, corresponding to the linear size of a magnitude $M \sim 6$ earthquake. Later analysis showed that the method was sensitive to $M \sim 5$. Within each box, a time series $N_i(t)$ is defined by counting how many earthquakes with magnitude greater than M_{\min} occurred during the time period t to $t + \delta t$. Next, the activity rate function $I_i(t_b, T)$ is defined as the average rate of occurrence of earthquakes in box i over the period t_b to T :

$$I_i(t_b, T) = \frac{\sum_{t=t_b}^T N_i(t)}{T - t_b}. \quad (\text{A1})$$

If t_b is held to be a fixed time, $I_i(t_b, T)$ can be interpreted as the i th component of a general, time-dependent vector evolving in an N -dimensional space (Tiampo et al., 2002b). Furthermore, it can be shown that this N -dimensional correlation space is defined by the eigenvectors of an $N \times N$ correlation matrix (Rundle et al., 2000a,b). The activity rate function is then normalized by subtracting the spatial mean over all boxes and scaling to give a unit-norm:

$$\hat{I}_i(t_b, T) = \frac{I_i(t_b, T) - \frac{1}{N} \sum_{j=1}^N I_j(t_b, T)}{\sqrt{\sum_{j=1}^N [I_j(t_b, T) - \frac{1}{N} \sum_{k=1}^N I_k(t_b, T)]^2}}. \quad (\text{A2})$$

The requirement that the rate functions have a constant norm helps remove random fluctuations from the system. Following the assumption of pure phase dynamics (Rundle et al., 2000a,b), the important changes in seismicity will be given by the change in the normalized activity rate function from the time period t_1 to t_2 :

$$\Delta \hat{I}_i(t_b, t_1, t_2) = \hat{I}_i(t_b, t_2) - \hat{I}_i(t_b, t_1). \quad (\text{A3})$$

This is simply a pure rotation of the N -dimensional unit vector $\hat{I}_i(t_b, T)$ through time. In order to both remove the last free parameter in the system, the choice of base year, as well as to further reduce random noise components, changes in the normalized activity rate function are then averaged over all possible base-time periods:

$$\Delta \hat{I}_i(t_0, t_1, t_2) = \frac{\sum_{t_b=t_0}^{t_1} \Delta \hat{I}_i(t_b, t_1, t_2)}{t_1 - t_0}. \quad (\text{A4})$$

Finally, the probability of change of activity in a given box is deduced from the square of its base averaged, mean normalized change in activity rate:

$$P_i(t_0, t_1, t_2) \propto [\Delta \hat{I}_i(t_0, t_1, t_2)]^2. \quad (\text{A5})$$

In phase dynamical systems, probabilities are related to the square of the associated vector phase function (Mori and Kuramoto, 1998; Rundle et al., 2000b). This probability function is often given relative to the background by subtracting off its spatial mean:

$$P'_i(t_0, t_1, t_2) \Rightarrow P_i(t_0, t_1, t_2) - \frac{1}{N} \sum_{j=1}^N P_j(t_0, t_1, t_2), \quad (\text{A6})$$

Where P' indicates the probability of change in activity is measured relative to the background.

Appendix B Updated method

1. The region of interest is divided into N_B square boxes with linear dimension Δx . Boxes are identified by a subscript i and are centered at x_i . For each box, there is a time series $N_i(t)$, which is the number of earthquakes per unit time at time t larger than the lower cut-off magnitude M_c . The time series in box i is defined between a base time t_b and the present time t .
2. All earthquakes in the immediate and eight surrounding regions of interest with magnitudes greater than a lower cutoff magnitude M_c are included. The lower cutoff magnitude M_c is specified in order to ensure completeness of the data through time, from an initial time t_0 to a final time t_2 . We use this extended, or Moore (1962), neighborhood to account for the uncertainty in event location and the arbitrary choice of where to center our boxes.
3. An intensity threshold is imposed such that only boxes that have historically been the most active are retained for analysis. The total number of earthquakes in each box from the initial time t_0 to the final time t_2 are counted, and boxes with counts less than the threshold are removed.
4. The seismic intensity in box i , $I_i(t_b, t)$, between two times $t_b < t$, can then be defined as the average number of earthquakes with magnitudes greater than M_c that occur in the box per unit time during the specified time interval t_b to t . Therefore, using discrete notation, we can write:

$$I_i(t_b, t) = \frac{1}{t - t_b} \sum_{t'=t_b}^t N_i(t'), \quad (\text{B1})$$

where the sum is performed over increments of the time series, which can be days to years.

5. Three time intervals are considered:
 - (a) A reference time interval from t_b to t_1 .
 - (b) A second time interval from t_b to t_2 , $t_2 > t_1$. The change interval over which seismic activity changes

are determined is then $t_2 - t_1$. The time t_b is chosen to lie between t_0 and t_1 . Typically we take $t_0=1932$, $t_1=1990$, and $t_2=2000$. The objective is to quantify anomalous seismic activity in the change interval t_1 to t_2 relative to the reference interval t_b to t_1 .

- (c) The forecast time interval t_2 to t_3 , for which the forecast is valid.

6. Our measure of anomalous seismicity in box i is the difference between the seismic intensities at t_1 and t_2 :

$$\Delta I_i(t_b, t_1, t_2) = I_i(t_b, t_2) - I_i(t_b, t_1). \quad (\text{B2})$$

7. In order to compare the intensities from different time intervals, we require that they have the same statistical properties. We therefore normalize each of the seismic intensities both individually over all choices for t_b and aggregately at each choice for t_b . This is performed by subtracting the mean seismic activity and dividing by the standard deviation of the seismic activity. The statistically normalized seismic intensity of box i is thus defined by

$$\begin{aligned} \Delta \tilde{I}_i(t_b, t_1, t_2) &= \frac{\Delta I_i(t_b, t_1, t_2) - \langle \Delta I_i(t_1, t_2) \rangle_T}{\sigma_T} \\ \Delta \hat{I}_i(t_b, t_1, t_2) &= \frac{\Delta \tilde{I}_i(t_b, t_1, t_2) - \langle \Delta \tilde{I}_i(t_b, t_1, t_2) \rangle_A}{\sigma_A}, \end{aligned} \quad (\text{B3})$$

where $\langle \Delta I_i(t_1, t_2) \rangle_T$ is the mean intensity difference of box i averaged over all choices of t_b , $\langle \Delta \tilde{I}_i(t_b, t_1, t_2) \rangle_A$ is the time averaged mean intensity difference averaged over all the boxes at each choice of t_b , and σ_T and σ_A are the respective standard deviations.

8. To reduce the relative importance of random fluctuations (noise) in seismic activity, we compute the average change in intensity, $\overline{\Delta I_i(t_0, t_1, t_2)}$ over many different pairs of normalized intensity maps having the same change interval:

$$\overline{\Delta I_i(t_0, t_1, t_2)} = \frac{1}{t_{max} - t_0} \sum_{t_b=t_0}^{t_{max}} \sqrt{\Delta \hat{I}_i(t_b, t_1, t_2)^2}, \quad (\text{B4})$$

where the sum is performed over increments of the time series for the distances between the normalized intensity differences and the background. In view of the fact that a time scale $\tau = t_2 - t_1$ has been implicitly chosen, the time t_{max} is chosen to be $t_{max} = t_1 - \tau$. This choice also gives the averaging time periods in the intervals t_b to t_1 and t_b to t_2 more equal weight, thereby excluding the possibility of large fluctuations caused by main shocks occurring just prior to t_1 .

9. We hypothesize that the probability of a future earthquake in box i , $P_i(t_0, t_1, t_2,)$, is proportional to the square of the mean intensity change:

$$P_i(t_0, t_1, t_2,) \propto \overline{\Delta I_i(t_0, t_1, t_2)^2}. \quad (\text{B5})$$

The constant of proportionality can be determined by requiring unit probability but is not important to the analysis.

10. To identify anomalous regions, we wish to compute the change in the probability $P_i(t_0, t_1, t_2,)$ relative to the background so that we subtract the mean probability over all boxes. We denote this change in the probability by

$$\Delta P_i(t_0, t_1, t_2) = P_i(t_0, t_1, t_2) - \langle P_i(t_0, t_1, t_2) \rangle_A, \quad (\text{B6})$$

where $\langle P_i(t_0, t_1, t_2) \rangle_A$ is the background probability for a large earthquake.

Hotspot pixels are defined to be the regions where $\Delta P_i(t_0, t_1, t_2)$ is positive. In these regions, $P_i(t_0, t_1, t_2)$ is larger than the average value for all boxes (the background level). Note that since the intensities are squared in defining probabilities, the hotspots may be due to either increases of seismic activity during the change time interval (activation) or due to decreases (quiescence). We hypothesize that earthquakes with magnitudes larger than M_c+2 will occur preferentially in hotspots during the forecast time interval t_2 to t_3 . A forecast map for all of California and its surrounding area using this procedure is given in Fig. C1. More details and results will be presented in a future publication.

Appendix C Previous research

The PI method was first introduced by Rundle et al. (2002) as an implication of the diffusive mean-field nature of earthquake dynamics. By treating seismicity as an example of a self-organizing threshold system they created forecast map for the occurrences of large earthquakes in southern California. At this time the method was known as Phase Dynamical Probability Change (PDPC).

Tiampo et al. (2002b) defined the PDPC method in mathematical terms and provided a rational explanation for each step of the process. They also performed likelihood tests against various null hypothesis and showed the PDPC method forecasts earthquakes better than random catalogs and better than a simple measure of past seismicity.

Holliday et al. (2005a) later performed a systematic analysis of the PDPC procedure. They varied the ordering of the steps and the parameter values and found optimal choices for the southern California region. The method at this time came to be known as Pattern Informatics (PI). Holliday et al. (2005b) then went on to investigate utilizing the PI method in a complex phase space. They determined that there is a small information gain for short-term (~ 5 year) forecasts when using complex eigenvectors rather than real-valued eigenvectors.

Nanjo et al. modified the PI method for use with the Japanese catalogs and successfully forecast the 23 October 2004 $M=6.8$ Niigata earthquake. Similarly, Chen and Holliday worked to modify the PI method for use with the Taiwanese catalogs. Insights from this work have led to the

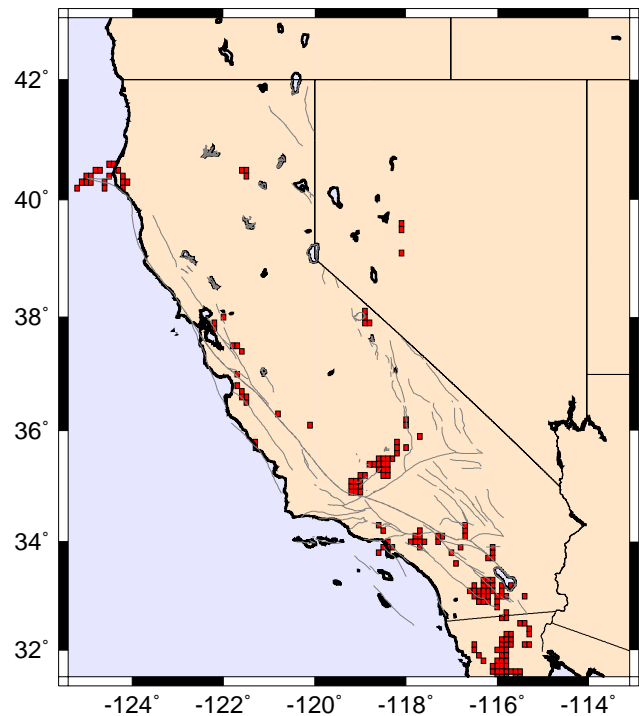


Fig. C1. Application of the modified PI method for all of California and its surrounding area. Colored areas are the forecast hotspots for the occurrence of $M \geq 5$ earthquakes during the period 2005–2015.

modified method for California forecasting presented in Appendix B and to a paper accepted for publication in *Geophysical Review Letters* (The 1999 Chi-Chi, Taiwan, earthquake as a typical example of seismic activation and quiescence).

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