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Paleomagnetism and U-Pb geochronology of the late Cretaceous Chisulryoung Volcanic Formation, Korea: tectonic evolution of the Korean Peninsula

Doohee Jeong¹, Yongjae Yu^{1*}, Seong-Jae Doh², Dongwoo Suk³ and Jeongmin Kim⁴**Abstract**

Late Cretaceous Chisulryoung Volcanic Formation (CVF) in southeastern Korea contains four ash-flow ignimbrite units (A1, A2, A3, and A4) and three intervening volcano-sedimentary layers (S1, S2, and S3). Reliable U-Pb ages obtained for zircons from the base and top of the CVF were 72.8 ± 1.7 Ma and 67.7 ± 2.1 Ma, respectively. Paleomagnetic analysis on pyroclastic units yielded mean magnetic directions and virtual geomagnetic poles (VGPs) as $D/I = 19.1^\circ/49.2^\circ$ ($\alpha_{95} = 4.2^\circ$, $k = 76.5$) and $VGP = 73.1^\circ\text{N}/232.1^\circ\text{E}$ ($A_{95} = 3.7^\circ$, $N = 3$) for A1, $D/I = 24.9^\circ/52.9^\circ$ ($\alpha_{95} = 5.9^\circ$, $k = 61.7$) and $VGP = 69.4^\circ\text{N}/217.3^\circ\text{E}$ ($A_{95} = 5.6^\circ$, $N = 11$) for A3, and $D/I = 10.9^\circ/50.1^\circ$ ($\alpha_{95} = 5.6^\circ$, $k = 38.6$) and $VGP = 79.8^\circ\text{N}/242.4^\circ\text{E}$ ($A_{95} = 5.0^\circ$, $N = 18$) for A4. Our best estimates of the paleopoles for A1, A3, and A4 are in remarkable agreement with the reference apparent polar wander path of China in late Cretaceous to early Paleogene, confirming that Korea has been rigidly attached to China (by implication to Eurasia) at least since the Cretaceous. The compiled paleomagnetic data of the Korean Peninsula suggest that the mode of clockwise rotations weakened since the mid-Jurassic. Such interesting variation of vertical rotations in the Korean Peninsula might result from the strike-slip motions of major faults developed in East Asia (the Tancheng-Lujiang fault to the northwest and the Korea-Taiwan strait fault to the southeast), near-field tectonic forcing of the subducting Pacific Plate beneath the Eurasian Plate, and far-field expressions of the India-Asia collision.

Background

East Asia is one of the world's poorly determined regions for reconstruction of Pangea (Besse and Courtillot 1991, 2002; Yang and Besse 2001; Enkin et al. 1992; Ma et al. 1993; Gilder et al. 2008; Pei et al. 2011). A robust reconstruction of Laurasia is possible only when Mesozoic paleogeography of the Korean Peninsula is well established, bridging the two sides of the Chinese cratons (e.g., Lee et al. 1987; Kim and Van der Voo 1990; Zhao et al. 1994, 1999; Doh et al. 1999, 2002; Park et al. 2003, 2005, 2007; Kim et al. 2009). The primary goal of paleomagnetism is to uncover information about the history of deformation and vertical rotations in geologic units. The present study is intended to analyze the variations in vertical axis block rotations in the Korean Peninsula with respect to more stable Chinese cratons.

The Korean Peninsula is located at the eastern end of the Eurasian Plate (Figure 1, inset), bounded by two major NNE-striking sinistral strike-slip fault systems of the Tancheng-Lujiang (Tan-Lu in abbreviation) fault to the northwest and the Korea-Taiwan strait fault to the southeast (Figure 1, inset). In particular, the Tan-Lu fault starts at eastern China and stretches north side of the Sino-Korean block over a distance of 1,000 km (Figure 1, inset). The southern part of the Korean Peninsula has five main tectonic units of Gyeonggi massif (GM), Ogcheon fold belt (OB), Ryongnam massif (RM), Gyeongsang Basin (GB), and Pohang Basin (PB) (e.g., Cluzel et al. 1991) (Figure 1). On the basis of combined paleomagnetic and paleontological analyses, it has been suggested that most tectonic units in Korea may have an affinity with South China while RM may have an affinity with North China (e.g., Kobayashi 1966, 1967; Doh and Piper 1994; Lee et al. 1996; Uno 1999; Uno and Chang 2000; Doh et al. 2002; Uno et al. 2004).

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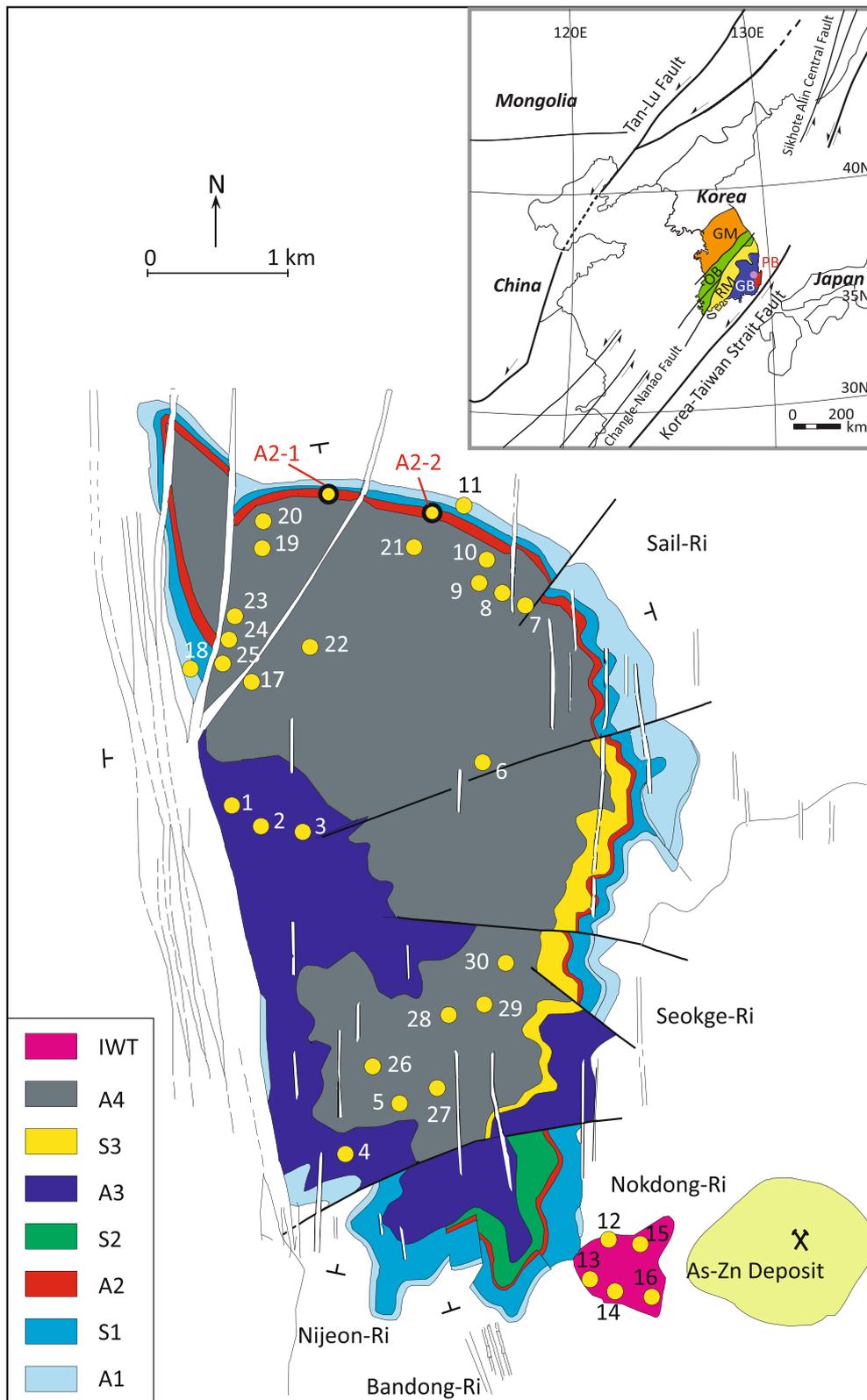


Figure 1 Location map of Korea (inset) and the Chisulyoung Volcanic Formation (CVF) outcrop pattern after Park (1990). The CVF has four ash-flow ignimbrite units (A1, A2, A3, and A4) and three intervening volcano-sedimentary layers (S1, S2, and S3). Yellow circles, locations of paleomagnetic sampling; GM, Gyeonggi massif; OB, Ogcheon fold belt; RM, Ryongnam massif; GB, Gyeongsang Basin; PB, Pohang Basin.

The present study deals with the paleomagnetism of late Cretaceous volcanic rocks in GB (Figure 1). Located in the southeastern part of the Korean Peninsula, GB is the largest non-marine sedimentary basin in the Sino-Korean block. Paleomagnetic investigation of GB is necessary to enhance our understanding of the Mesozoic tectonic evolution of the Sino-Korean block. In the present study, we report paleomagnetic data along with high-precision zircon U-Pb dating from the Chisulryoung Volcanic Formation (CVF) in GB, which can shed some light on the late Cretaceous tectonism in Korea, as well as possible tectonic correlation with the Eurasian Plate.

Methods

Geology

The CVF of southeastern Korea forms an extrusive outlier (6 × 3 km) around Chisulryoung Mountain (765 m) (pink circle in Figure 1 inset). According to Reedman et al. (1987), the CVF was formed by two sequential cycles of subaerial explosive silicic volcanic activities in which an emplacement of thick ash-flows intercalated with volcanoclastic sedimentary rocks. It is typically a basin, circular in form whose underlying sedimentary rocks inclined with centripetal dips (Figure 1). Overall, the CVF lies within a broad structural basin regarded as a down sag caldera (Reedman et al. 1989).

The CVF comprises mainly densely to weakly welded acidic ash-flow tuffs which have been divided into four members (A1, A2, A3, and A4) and three intervening volcano-sedimentary members (S1, S2, and S3) according to their crystal content, grain-size distribution of pyroclastic material, and degree of consolidation and welding (Park and Kim 1985a, b; Park 1990) (Figure 2). As a first major eruption cycle, a sequence of lower members (A1, S1, A2, S2, and A3) predates the intrusion of the Cheogwari granodiorite whose hornblende K-Ar age is 76.3 ± 11.7 Ma (Reedman et al. 1989). The A4 member is a product of the second eruption cycle, and the S3 member represents sedimentations that occurred between the two major eruption cycles (Reedman et al. 1989; Park 1990). To put it differently, the two major eruption cycles were separated by the intrusion of the Cheogwari granodiorite (Reedman et al. 1989). In particular, the A4 member comprises approximately 100 to 120 m of loosely welded pumice-rich tuff (A4a) overlain by about 80 to 120 m of densely welded high-density lithic-rich tuff (A4b) (Figure 2). The second eruptive cycle was terminated with the intrusion of the Gadaeri granite whose biotite K-Ar age is 63.9 ± 1.8 Ma (Reedman et al. 1989) (Figure 2). An intrusive welded tuff (IWT) plug, with an elliptical cross section (700 × 400 m), has a eutaxitic texture nearly parallel to its subvertical contacts.

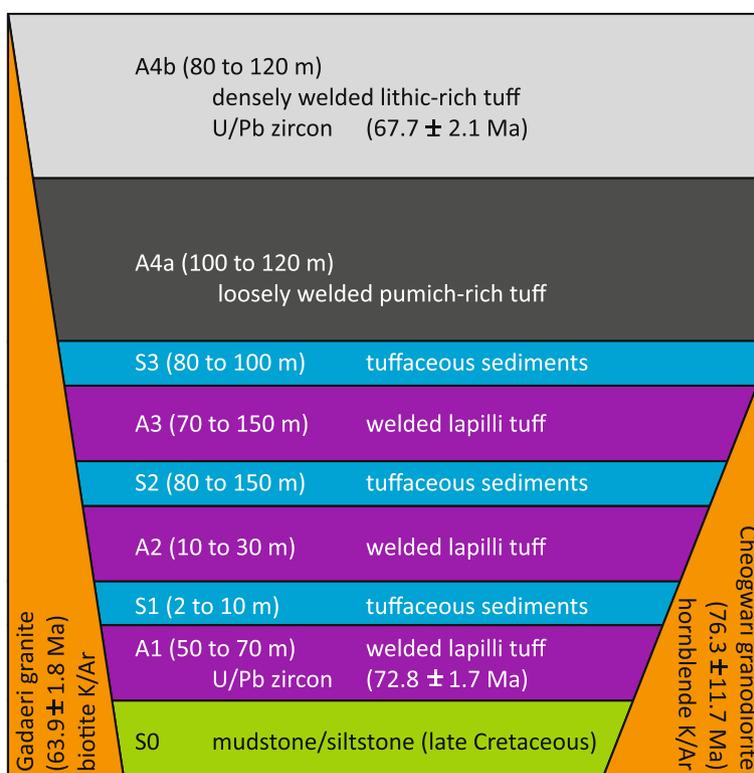


Figure 2 Schematic section illustrating relative stratigraphic positions of the CVF, lithologies, tentative thickness, and available age determinations.

The IWT cuts; hence, it post-dates the Cheogwari granodiorite. Indeed, the IWT yielded sericite K-Ar age of 65.1 ± 1.8 Ma, slightly older than the emplacement age of the Gadaeri granite (Reedman et al. 1987, 1989; Park 1990).

Sampling and laboratory procedures

Rheomorphic ash-flow tuffs occur as hot lateral flows invade the pumice deposits, producing extreme flattening and stretching of dense ignimbrites. Whenever possible, we avoided rheomorphic ignimbrites as they might compromise reliable remanence record. Fifty-four oriented samples were taken from six sites by *in situ* drilling. These sites are distributed in the western part of the study area where slopes are relatively gentle. In 24 sites where drilling was inapplicable due to high topography and steep slopes along the mountain side, oriented block samples were collected. Six to ten 2.5-cm cylindrical specimens of each sample were alternating-field (AF) demagnetized in 12 steps to 90 mT using a Molspin AF demagnetizer (ASC Scientific, Carlsbad, CA, USA). In addition, other specimens were thermally demagnetized, usually in 50°C steps to 500°C and then in smaller steps to 700°C using a non-inductive thermal demagnetizer (ASC Scientific Model TD-48, ASC Scientific, Carlsbad, CA, USA). Optimum demagnetization step to calculate a mean remanence direction was defined as that producing the minimum dispersion in direction. Sample storage, demagnetization, and spinner magnetometer measurements were performed in a magnetically shielded space with a nominal ambient field of <200 nT at the Paleomagnetism Laboratory, Korea University.

To date, the relative age sequence of the CVF was indirectly established on the basis of the K-Ar ages determined from the intervening intrusive granitoids (Figure 2). In an attempt to directly determine the age of the pyroclastic units, high-precision age dating was carried out. At first, zircons from the A1, A4, and IWT were dated by the U-Pb method using the Sensitive High Resolution Ion MicroProbe (SHRIMP) at the Korea Basic Science Institute. Details of operation and data reduction protocols were similar to those described by Yi et al. (2012). Zircons were extracted from crushed and pulverized samples, using standard magnetic and density separation. Cathodoluminescence (CL) and back-scattered electron images were obtained using a scanning electron microscope (SEM) of JEOL 6610LV (JEOL Ltd., Akishima-shi, Japan). Zircon inclusions were identified by using an INCA x-act energy-dispersive spectrometer (Oxford Instruments, Abingdon, UK). Concentrations of U and Th were calculated with reference to SL13 (U = 238 ppm). The measured ratio of $^{206}\text{Pb}/^{238}\text{U}$ was calibrated for the standard of 1.1-Gyr-old FC1 zircon (Paces and Miller

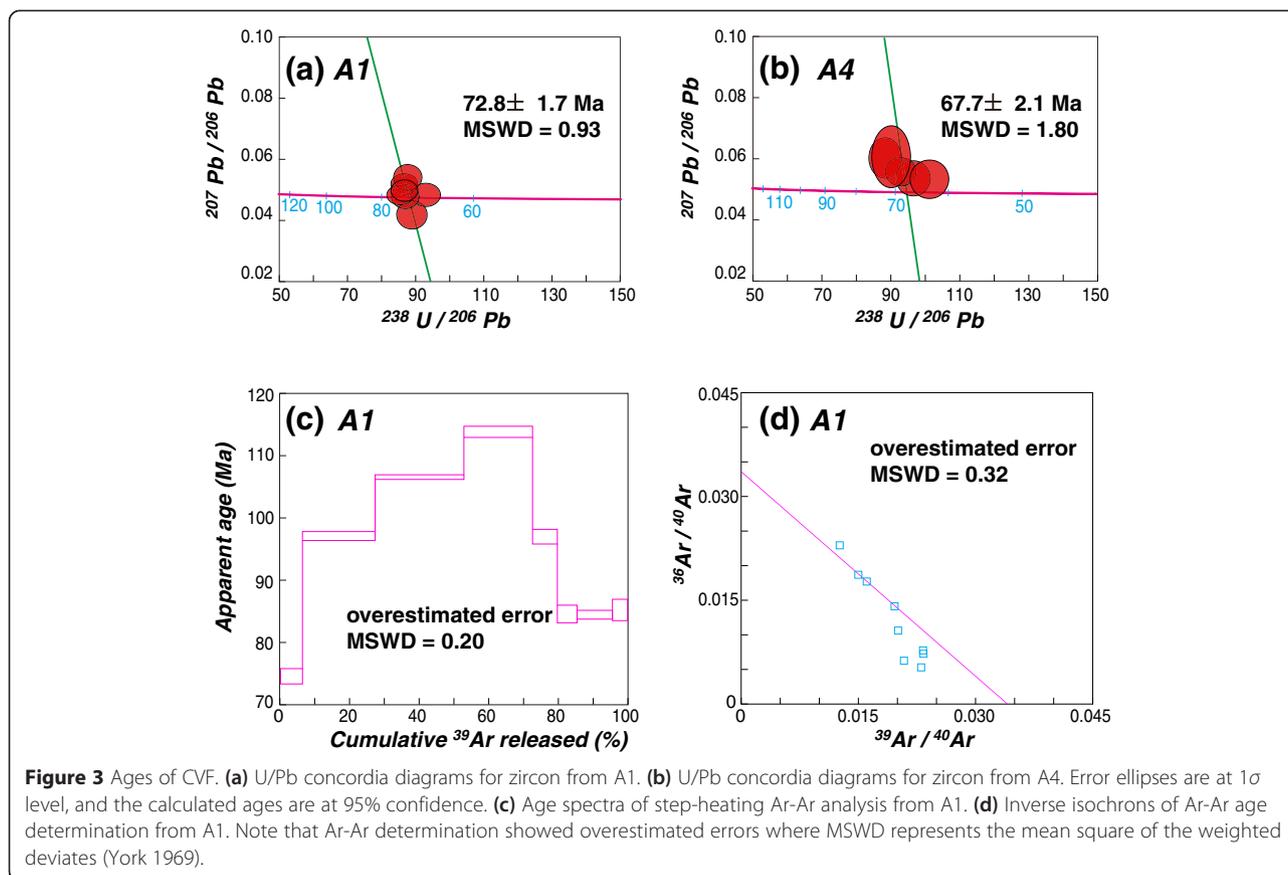
1993). Concordia diagrams and age determination were made after excluding the outliers on the basis of the *t*-test using two commonly used programs of Squid 2.50 and Isoplot 3.71 (Ludwig 2008, 2009). Second, high-precision $^{40}\text{Ar}/^{39}\text{Ar}$ step-heating analyses were performed at the Argon Geochronology Laboratory, Oregon State University using the freshest matrix crushed from rock samples.

Seven zircon grains of A1 member yielded concordant U-Pb ages of 72.8 ± 1.7 Ma with a mean square of the weighted deviates (MSWD) of 0.93 (Figure 3a). Similarly, eight zircon grains of A4 member displayed concordant U-Pb ages of 67.7 ± 2.1 Ma with a MSWD (York 1969) of 1.80 (Figure 3b). However, we failed to obtain reliable concordant U-Pb ages from the IWT because of enormously high concentrations of uranium. We were also unable to retrieve scientifically meaningful $^{40}\text{Ar}/^{39}\text{Ar}$ ages for the rocks in the CVF. For instance, both step-heating plateau age spectra (Figure 3c) and inverse isochron regression fittings (Figure 3d) are all unacceptable with overestimated errors. Despite limited success, U-Pb zircon ages of A1 (72.8 ± 1.7 Ma) and A4 (67.7 ± 2.1 Ma) confirm that the two major eruption cycles of the CVF were separated by the intrusion of the Cheogwari granodiorite (hornblende K-Ar age of 76.3 ± 11.7 Ma) and the magmatic activity of the CVF was terminated with the intrusion of the Gadaeri granite (biotite K-Ar age of 63.9 ± 1.8 Ma).

Results

Hysteresis loops were measured from 44 rock chips on an alternating gradient force magnetometer at the University of Toronto, Canada. Values of saturation magnetization (M_s), saturation remanence (M_r), and coercive force (B_c) were determined from hysteresis loops. Hysteresis parameters were determined after appropriate paramagnetic slope correction. Representative hysteresis loops are given in Figure 4. In most samples, hysteresis loops for fine- to coarse-grained magnetite are typically observed (Figure 4a,b). For example, sample A4-30-4 shows a high remanence ratio (M_r/M_s) of ≈ 0.4 , similar to near single-domain (SD) or fine pseudo-single-domain (PSD) characteristics (Figure 4a). On the other hand, A4-17-5 shows more multidomain (MD)-like behavior of low remanence ratio of ≈ 0.2 (Figure 4b). High coercivities over 100 mT and high remanence ratio over 0.5 from the four chips of site 27 suggest the presence of highly coercive mineral other than magnetite (Figure 4c). In a Néel (1955) plot of remanence ratio (M_r/M_s) versus coercivity (B_c), a total of 40 representative magnetite-containing chips follow the trend of the SD-MD mixing lines (Figure 4d) according to the criteria of Tauxe et al. (2002).

Demagnetization results were displayed in geographic coordinates in terms of total intensity variation of magnetization, equal-area stereographic plot, and vector



projections using a commonly used PuffinPlot (Lurcock and Wilson 2012). The characteristic remanent magnetization (ChRM) directions were calculated from a principal component analysis (Kirschvink 1980). Virtual geomagnetic poles (VGPs) were calculated for each site, and mean paleomagnetic poles and their associated confidence limits were accordingly defined. Directions of natural remanent magnetization (NRM) prior to demagnetization treatments were slightly dispersed for each site, suggesting a weak viscous magnetic overprint. Magnetically weak samples yielded either an aberrant direction or a direction that progressively approaches the characteristic direction but fails to achieve a stable end point. These aberrant samples (see Table 1) were not used in mean calculation. In a few sites, a great circle method (McFadden and McElhinny 1988) was applied in mean calculation whenever NRM directions follow a great circle demagnetization trajectory.

The first eruptive cycle (A1, A2, and A3)

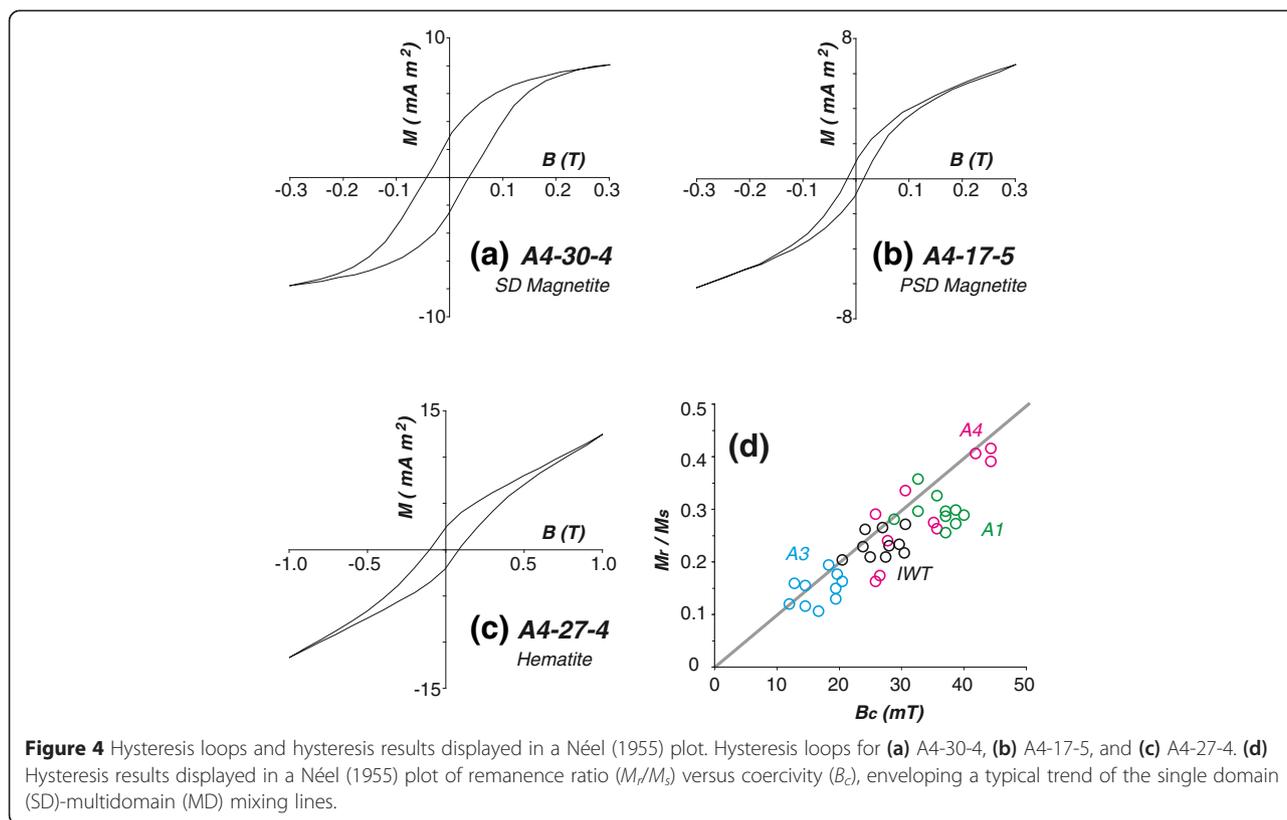
Forty-six oriented samples were collected at six sites from the first eruptive cycle: two (sites 11 and 18) in A1 and four (sites 1, 2, 3, and 4) in A3 (Figures 1 and 2). A weak spurious magnetization is removed usually <10 mT or 150°C, after which a stable ChRM is identified

from 10 to 90 mT (Figure 5a) or 100°C to 580°C (Figure 5b). The maximum unblocking temperature of 580°C reveals magnetite as a main remanence carrier (Figure 5b). The stable components are uniformly of normal polarity both for A1 and A3 (Figure 5, Table 1).

In A3, one of the four sites (site 4) has an anomalous but well-clustered directions (Table 1). It is unclear whether site 4 records an isolated yet unidentified block movement or a short-lived geomagnetic polarity excursion. Nonetheless, we discarded the results of site 4 in mean calculation (Table 1). Site mean directions and corresponding paleopoles are $D/I = 19.1^\circ/49.2^\circ$ ($\alpha_{95} = 4.2^\circ$, $k = 76.5$) and $VGP = 73.1^\circ\text{N}/232.1^\circ\text{E}$ ($A_{95} = 3.7^\circ$, $N = 3$) for A3.

The mean results for A1 were based on two sites ($N = 2$; sites 11 and 18). As the Fisher site mean determinations require at least three (i.e., $N \geq 3$) individual paleomagnetic mean directions, we used combined specimen mean directions ($N = 11$; $N = 4$ for site 11 and $N = 7$ for site 18). Combined mean directions and corresponding paleopoles are $D/I = 24.9^\circ/52.9^\circ$ ($\alpha_{95} = 5.9^\circ$, $k = 61.7$) and $VGP = 69.4^\circ\text{N}/217.3^\circ\text{E}$ ($A_{95} = 5.6^\circ$, $N = 11$) for A3 (Table 1).

For A2, 14 samples were collected at two sites (A2-1 and A2-2 in Figure 1). Unfortunately, all the samples were weakly magnetized and become fractured during



thermal demagnetization. Hence, we were unable to retrieve reliable paleomagnetic results from A2.

Intrusive welded tuff

Forty-four oriented samples were collected at five sites (sites 12, 13, 14, 15, and 16) from the IWT. All the samples exhibited stable demagnetization behavior (Figure 6). During AF demagnetization, the median destructive field (MDF) was over 30 to 40 mT (Figure 6a). Thermal demagnetization data show narrow unblocking temperature spectra of 500°C to 600°C, consistent with a fine-grained magnetite as the

principal carrier of magnetic remanence (Figure 6b). Paleomagnetic mean directions and paleopoles for IWT are $D/I = 285.7^\circ/82.3^\circ$ ($\alpha_{95} = 5.3^\circ$, $k = 17.6$) and VGP = $38.1^\circ\text{N}/110.5^\circ\text{E}$ ($A_{95} = 10.1^\circ$, $N = 5$) (Table 2).

Paleomagnetic data from the IWT showed high magnetic inclination and radially scattered magnetic declination (Table 2). Rheomorphic ash-flow tuffs develop preferred alignment of pyroclasts that occurred shortly after the emplacement of laminar viscous flow (Wolff and Wright 1981). Then, it is likely that high magnetic inclination of IWT results from rheorphism (Uno et al. 2013).

Table 1 Paleomagnetic results in the first eruption cycle

Member	Site	φ_s	λ_s	n_1/n_2	D_i	I_i	k	α_{95}	φ_p	λ_p	dp	dm
A3	1	129.14	35.40	6/10	16.2	48.0	87.8	7.2	239.2	74.9	6.2	9.4
	2	129.14	35.40	5/7	13.8	49.3	59.6	10.0	239.5	77.3	8.8	13.3
	3	129.14	35.40	5/7	23.2	50.3	65.4	9.5	225.2	70.2	8.5	12.7
	4 ^a	129.14	35.39	9/9	134.7	15.0	28.9	9.7	183.1	-29.4	5.1	10.0
	Mean				19.1	49.2	76.5	4.2	232.1	73.1	$A_{95} =$	3.7
A1	11	129.15	35.41	4/8	16.6	52.2	82.2	10.2	225.1	76.0	9.6	14.0
	18	129.14	35.41	7/11	29.8	53.1	59.7	7.9	214.5	65.5	7.6	11.0
	Mean				24.9	52.9	61.7	5.9	217.3	69.4	$A_{95} =$	5.6

φ_p/λ_p , sampling site longitude/latitude; n_1/n_2 , number of samples used in calculating means/demagnetized at each site; D_i/I_i , mean declination/inclination; k , precision parameter; α_{95} , 95% confidence limit; φ_p/λ_p , VGP longitude/latitude; dp/dm , confidence limits; A_{95} , the angular radius of an ellipse within which the mean pole lies within 95% confidence. ^aNot included in mean calculation.

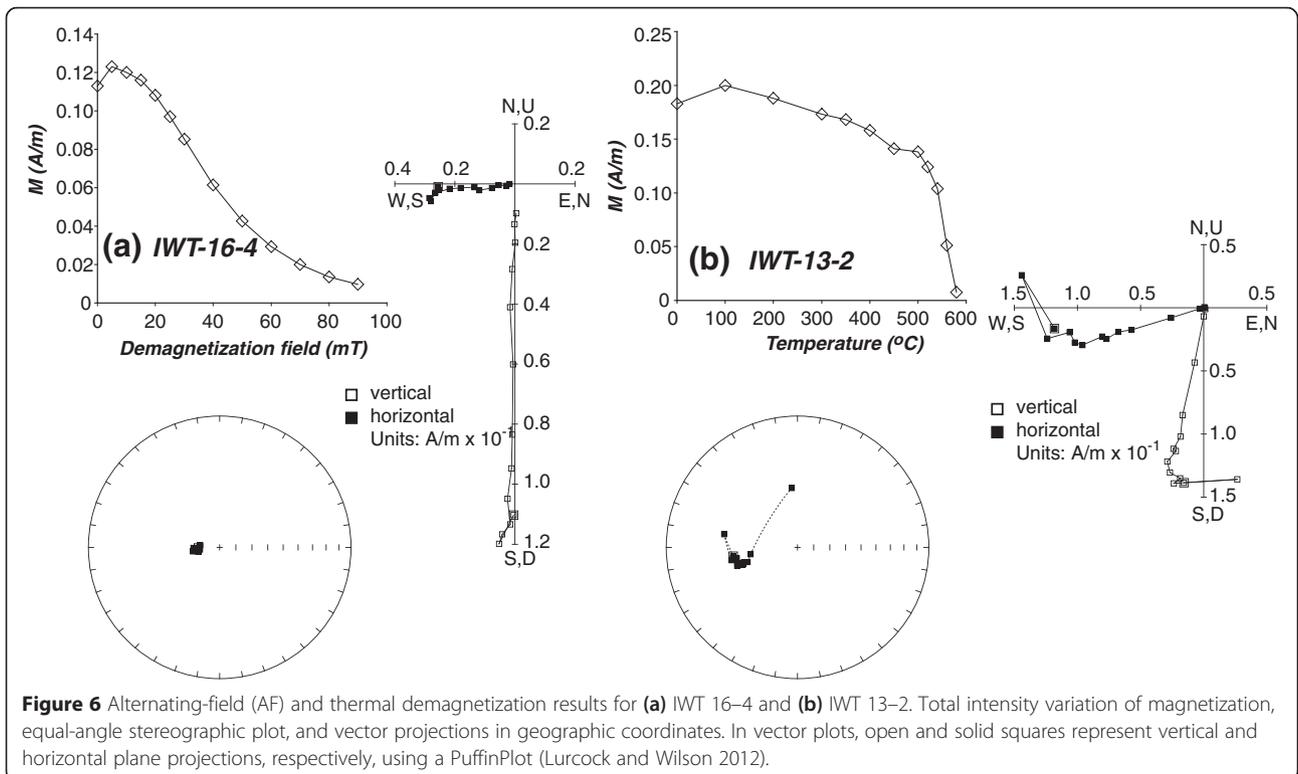
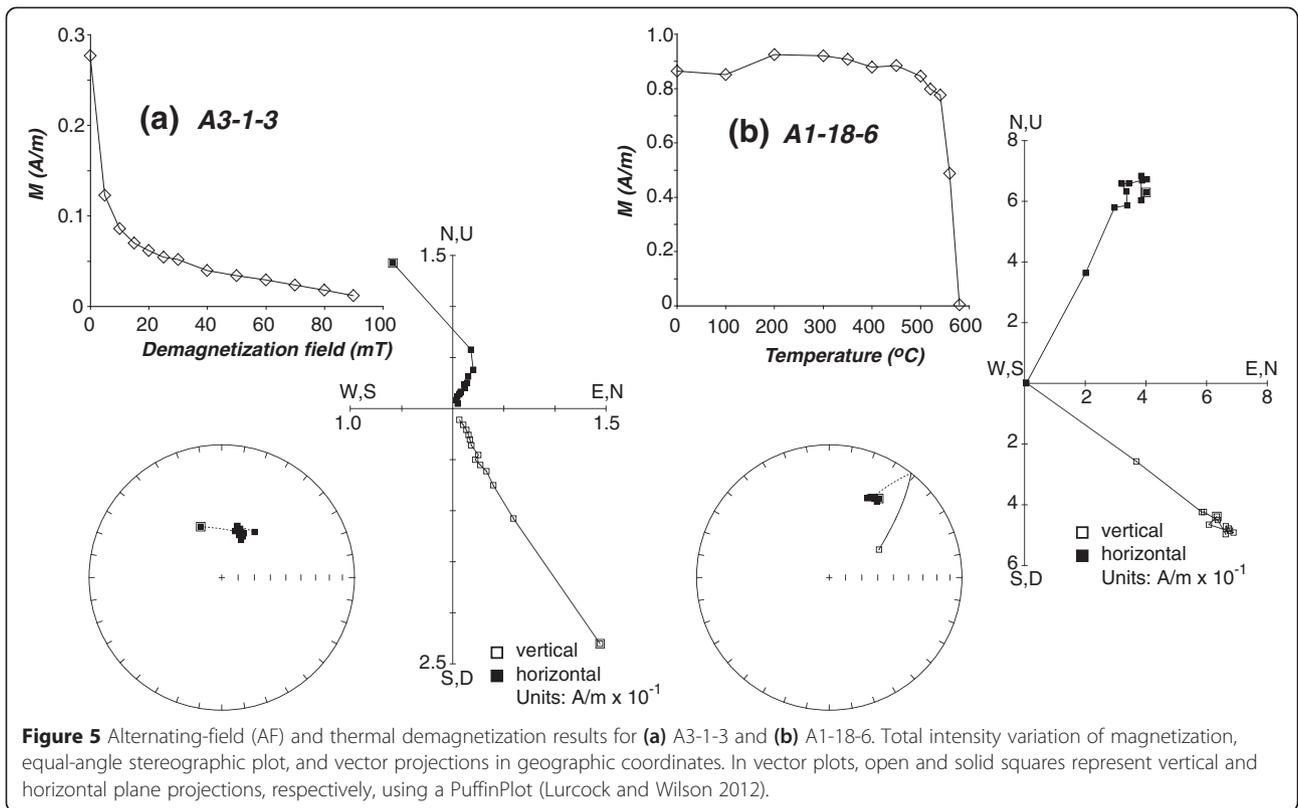


Table 2 Paleomagnetic results in the second eruption cycle

Member	Site	φ_s	λ_s	n_1/n_2	D_i	I_i	K	α_{95}	φ_p	λ_p	dp	dm	
A4	5 ^a	129.15	35.40	4/6	138.8	-54.7	24.9	18.8	263.7	-56.8	27.4	32.1	
	6	129.15	35.40	4/5	17.5	13.7	27.1	18.0	275.6	57.3	9.4	18.4	
	7	129.15	35.41	6/7	27.2	51.0	24.8	11.5	220.7	67.1	10.5	15.6	
	8	129.15	35.41	6/7	36.5	33.0	60.5	8.7	237.8	53.3	5.6	9.9	
	9	129.15	35.41	5/6	36.1	21.3	35.7	13.0	246.9	49.2	7.2	13.7	
	10	129.15	35.41	4/6	20.5	23.4	24.4	19.0	265.4	60.3	10.8	20.2	
	17	129.14	35.41	11/11	346.1	48.9	104.7	4.5	17.5	77.0	3.9	5.9	
	19	129.14	35.41	14/18	346.2	36.5	15.5	10.4	351.6	70.6	7.1	12.1	
	20	129.14	35.41	17/23	358.2	32.3	37.6	5.9	314.7	72.1	3.8	6.7	
	21	129.15	35.41	24/24	16.4	36.2	36.0	5.0	261.5	69.0	3.4	5.8	
	22	129.14	35.41	14/14	359.4	41.3	16.6	10.1	311.9	78.3	7.5	12.3	
	23	129.14	35.41	9/21	201.3	-41.6	12.2	15.4	65.3	-68.3	11.5	18.8	
	24	129.14	35.41	10/16	192.2	-53.0	8.3	17.8	46.0	-79.8	17.1	24.7	
	25	129.14	35.41	10/13	357.0	48.4	25.5	9.8	332.8	83.5	8.4	12.9	
	26	129.15	35.39	13/14	186.6	-46.9	10.2	13.7	89.7	-80.8	11.4	17.7	
	27	129.15	35.39	19/22	192.2	-55.3	15.4	8.8	33.0	-80.1	8.9	12.5	
	28	129.15	35.39	14/14	11.1	60.7	22.1	8.6	180.0	79.3	10.0	13.1	
	29	129.15	35.39	8/10	3.9	32.3	23.5	11.7	297.2	71.8	7.4	13.2	
	30	129.15	35.40	17/19	2.7	30.9	30.3	6.6	301.1	71.1	4.1	7.4	
	Mean	Normal				10.7	50.0	31.1	7.2	243.4	79.9	$A_{95}=\text{---}$	6.4
	Mean	Reversed				192.3	-39.7	112.7	8.7	266.0	-73.3	$A_{95}=\text{---}$	6.3
	Mean					10.9	50.1	38.6	5.6	242.4	79.8	$A_{95}=\text{---}$	5.0
	IWT	12	129.16	35.39	9/9	99.9	84.2	56.3	6.9	142.6	32.7	13.4	13.6
		13	129.16	35.39	10/10	240.4	70.6	12.7	14.1	98.1	14.0	21.2	24.4
		14	129.16	35.38	10/10	303.2	76.3	27.5	9.4	97.5	45.7	16.1	17.4
		15	129.16	35.38	8/8	338.4	73.8	22.6	11.9	106.1	61.8	19.3	21.4
		16	129.16	35.38	7/7	265.5	83.0	47.9	8.8	112.7	33.2	16.8	17.2
Mean					285.7	82.3	17.6	5.3	110.5	38.1	$A_{95}=\text{---}$	10.1	

φ_p/λ_p , sampling site longitude/latitude; n_1/n_2 , number of samples used in calculating means/demagnetized at each site; D_i/I_i , mean declination/inclination; K , precision parameter; α_{95} , 95% confidence limit; φ_p/λ_p , VGP longitude/latitude; dp/dm , confidence limits; A_{95} , the angular radius of an ellipse within which the mean pole lies within 95% confidence. ^aNot included in mean calculation.

The second eruptive cycle (A4)

A total of 242 oriented samples were collected at 19 sites (sites 5, 6, 7, 8, 9, 10, 17, 19, 20, 21, 22, 23, 24, 25, 26, 27, 28, 29, and 30) from the second eruptive cycle. Progressive AF and thermal demagnetization isolated a well-defined stable magnetization (Figure 7). An initial plateau with relatively high MDF in AF demagnetization (Figure 7a) and a sharp drop of magnetization near 580°C in thermal demagnetization (Figure 7b) are characteristics of fine-grained magnetite.

A unique demagnetization behavior was observed from the seven samples in site 27 (Figure 8). These samples are completely resistant to AF demagnetization (Figure 8a) and have maximum unblocking temperatures approximately 680°C (Figure 8b), suggesting that the remanence is carried by hematite. The mean direction of

these seven hematite-bearing samples is consistent with that of twelve other magnetite-bearing samples from the same sampling site in 95% confidence limits. It is difficult to envisage a low-temperature chemical process that would generate both magnetite and hematite. It is possible that a high-temperature deuteric oxidation formed both magnetite and hematite (e.g., Schlinger et al. 1991) during the early stage of emplacement. If so, the NRM in site 27 is a combination of thermoremanent magnetization (TRM) carried by magnetite and thermochemical remanent magnetization (TCRM) acquired during high-temperature oxidation of magnetite to hematite (e.g., May et al. 1986).

In A4, the ChRM directions show both normal and reversed polarities (Table 2). For instance, the ChRM directions of four sites (sites 23, 24, 26, and 27) are

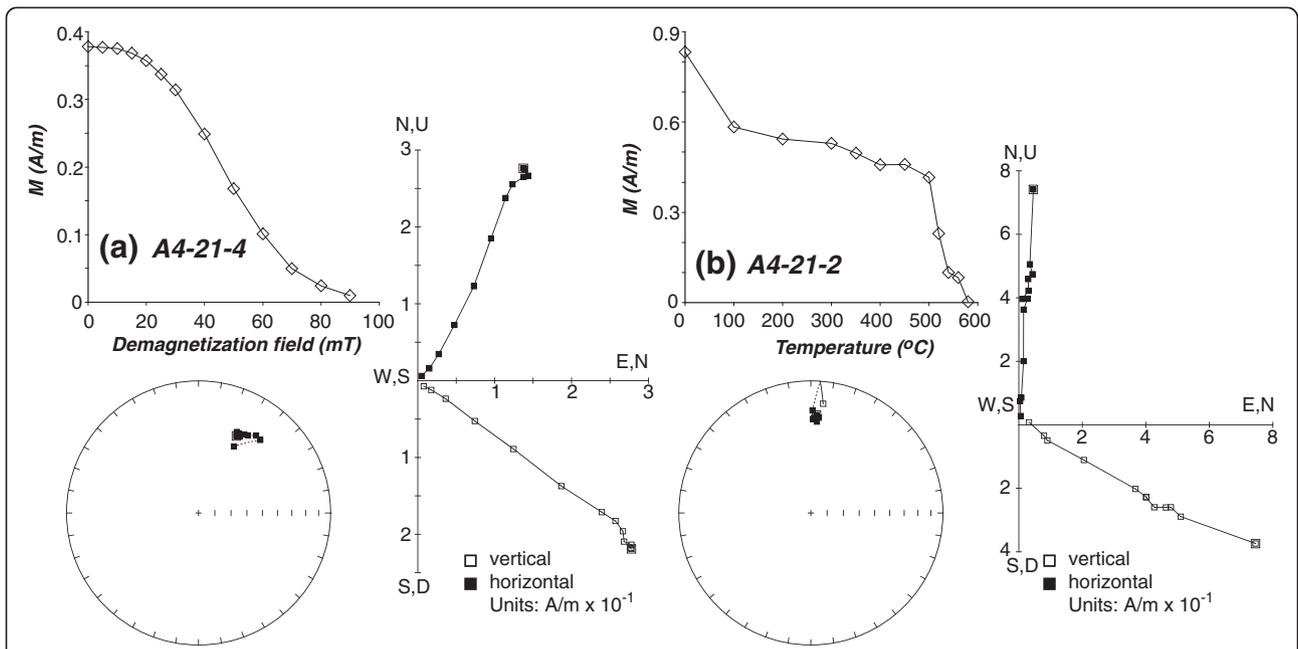


Figure 7 Alternating-field (AF) and thermal demagnetization results for (a) A4-21-4 and (b) A4-21-2. Total intensity variation of magnetization, equal-angle stereographic plot, and vector projections in geographic coordinates. In vector plots, open and solid squares represent vertical and horizontal plane projections, respectively, using a PuffinPlot (Lurcock and Wilson 2012).

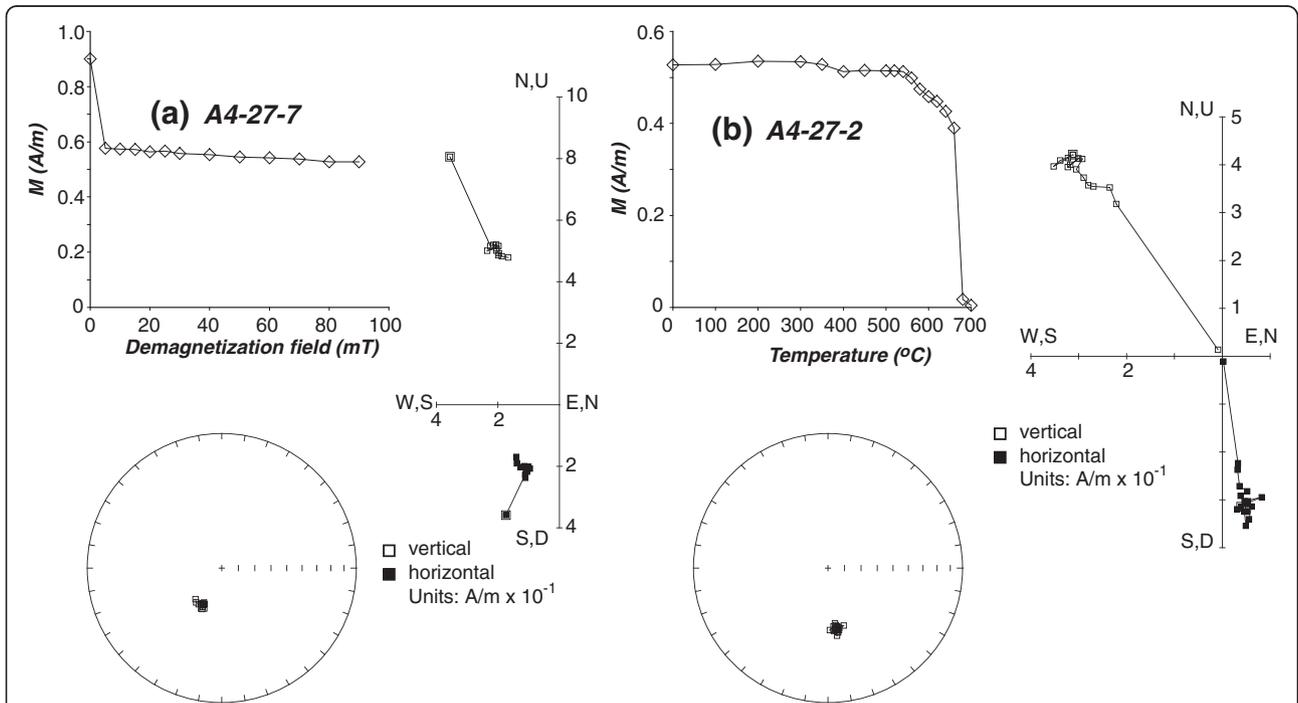


Figure 8 An example of (a) AF and (b) thermal demagnetization of NRM for A4 in geographic coordinates. NRM showed the presence of a component which is resistant AF up to 90 mT and heating up to 680°C. In vector plots, open and solid squares represent vertical and horizontal plane projections, respectively, using PuffinPlot (Lurcock and Wilson 2012).

reversely magnetized, while those of fourteen sites (sites 6, 7, 8, 9, 10, 17, 19, 20, 21, 22, 25, 28, 29, and 30) are normal (Table 2). Paleomagnetic site mean directions and paleopoles are $D/I = 10.7^\circ/50.0^\circ$ ($\alpha_{95} = 7.2^\circ$, $k = 31.1$) and $VGP = 79.9^\circ\text{N}/243.4^\circ\text{E}$ ($A_{95} = 6.4^\circ$, $N = 14$) for normal polarity and $D/I = 192.3^\circ/-39.7^\circ$ ($\alpha_{95} = 8.7^\circ$, $k = 112.7$) and $VGP = 73.3^\circ\text{S}/266.0^\circ\text{E}$ ($A_{95} = 6.3^\circ$, $N = 4$) for reversed polarity (Table 2). The mean ChRM directions from fourteen normal polarity sites are anti-parallel to those of four reversed polarity sites (Table 2). Indeed, the ChRM directions pass the reversal test for paleomagnetic stability (Figure 9). When combined, the paleomagnetic site mean direction and paleopole for the entire A4 are $D/I = 10.9^\circ/50.1^\circ$ ($\alpha_{95} = 5.6^\circ$, $k = 38.6$) and $VGP = 79.8^\circ\text{N}/242.4^\circ\text{E}$ ($A_{95} = 5.0^\circ$, $N = 18$) (Table 2).

Discussion

High-precision U-Pb zircon age obtained from the base of the CVF (i.e., A1) is 72.8 ± 1.7 Ma (Figures 2 and 3). U-Pb zircon geochronology is unavailable for the overlying sequential tuff layers and tuffaceous sediments for the first eruption cycle. On the basis of normal polarities observed for both A1 and A3, the first eruption cycle of the CVF can be constrained as chron 32N in late Campanian. On the other hand, U-Pb zircon age of 67.7 ± 2.1 Ma was obtained from the stratigraphically higher A4 (Figure 2). While magnetite-bearing tuffs retained TRM of normal polarity, hematite-containing tuffs preserved TCRM of reversed polarity (Figure 8, Table 2). Existence of dual polarities in the same site implies that the second eruption cycle of the CVF may record polarity transition from chron C31N to C30R in late Maastrichtian. Overall, the CVF reflects late Cretaceous volcanism that definitely predates the Paleogene (Figure 2).

Paleomagnetic stability of pyroclastics is influenced by the structural correction to restore the beds to horizontal, degree of bedding tilts, and the caldera affinity. In

the present study, no structural correction was employed in analyzing data. Determining systematic variation of attitudes of the intervening volcanic-sedimentary layers was unavailable because the horizontal continuity of A1, S1, A2, S2, A3, and S3 was poor. Instead, we relied on the between- or among-site directional consistency of paleomagnetic directions as a sideline evidence in assessing the quality of paleomagnetic directions. For instance, it is obvious that paleomagnetic mean direction is strongly influenced by the caldera affinity as IWT showed high inclination (Table 2). Therefore, the paleomagnetic mean direction of IWT was excluded in further tectonic discussion. On the other hand, results from A4 suggest that paleomagnetic directions were less relevant to the distances from sampling sites to volcanic-sedimentary boundaries or local faults (Tables 1 and 2).

Due to their inherent sporadic nature, pyroclastic rocks fossilize spot-readings of geomagnetic field vector during volcanic eruptions. As CVF does not reflect massive ever-lasting eruptions, it is difficult to imagine that secular variation is completely averaged out. However, the presence of dual polarities in flow A4 provides a strong rationale for the averaging out of secular variation (Figure 9).

The paleomagnetic poles of A1, A3, and A4 show a counterclockwise swing in the third quadrant (Figure 10). The apparent polar wander path (APWP) of CVF covers almost 60° swing of VGP longitude (from eastern Alaska to northern Greenland) but about 10° increase of VGP latitude (Figure 10). VGP positions of A1 and A3 fell squarely well on the paleopoles of China in Cretaceous age (Gilder and Courtillot 1997; Huang et al. 2007). In particular, the paleopole of A3 is statistically identical to a recently refined late Cretaceous reference pole for the stable South China (Li et al. 2013). On the other hand, the paleopole of A4 is statistically indistinguishable to the paleopoles of China in early Tertiary (Zhao et al.

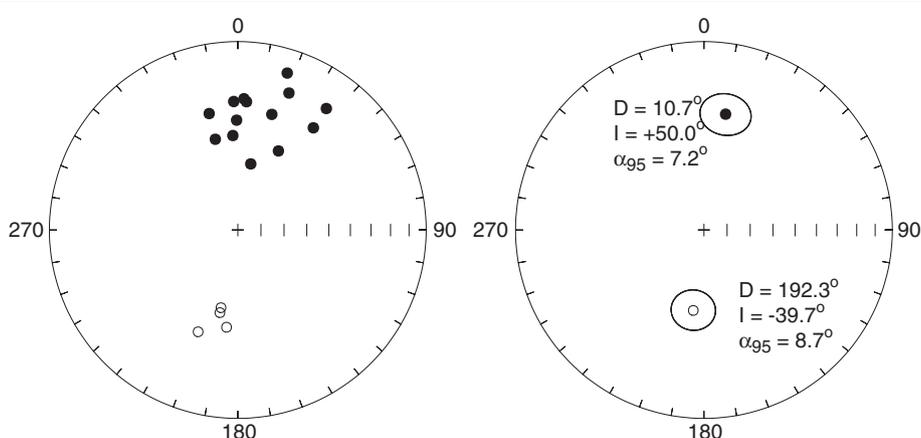
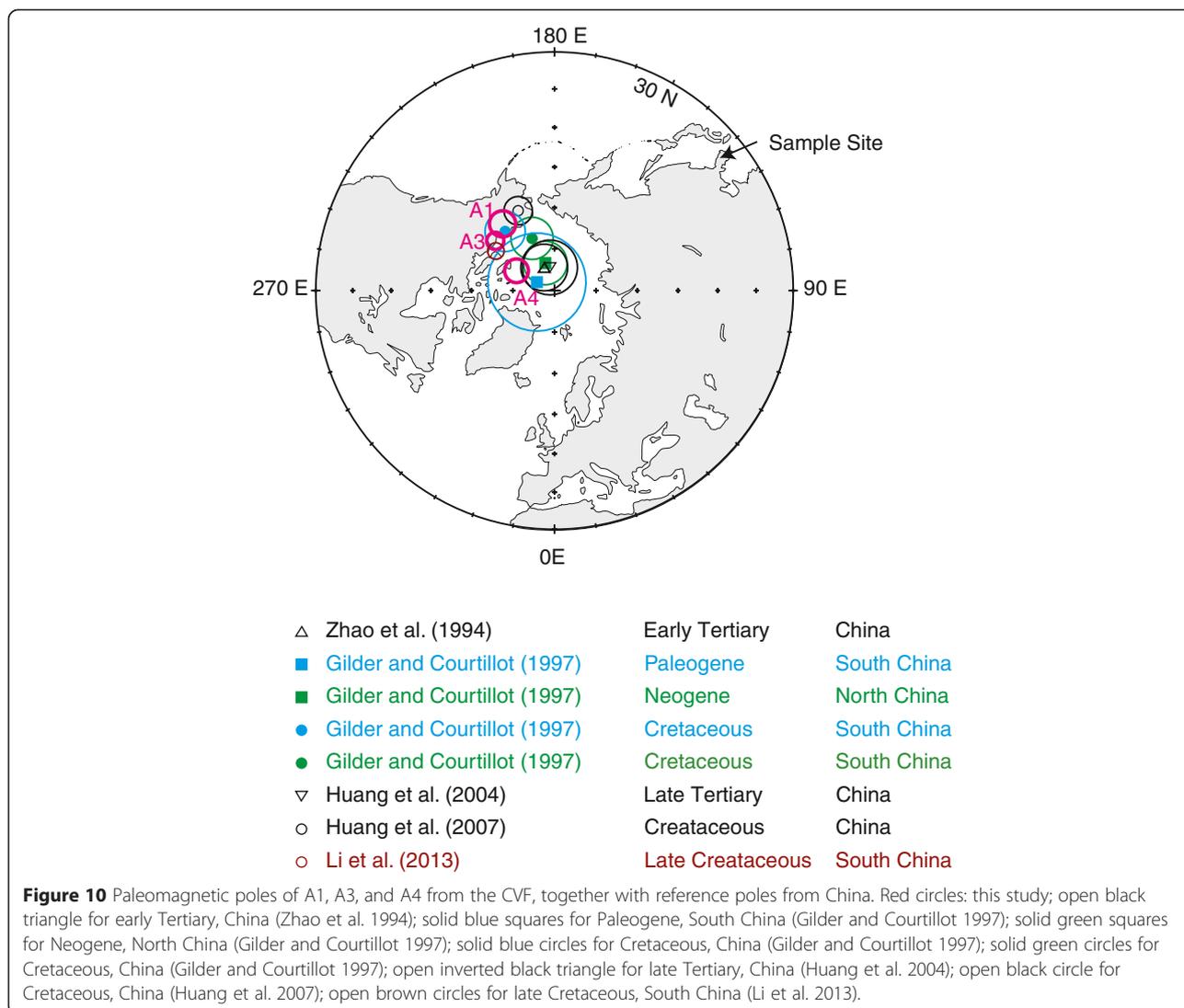


Figure 9 Paleomagnetic site mean directions of A4 pass the reversal test.



1994), Paleogene to Neogene (Gilder and Courtillot 1997), and/or late Tertiary (Huang et al. 2004). The Cretaceous paleopoles of North China Block (NCB) and South China Block (SCB) are the result of averaging both Early and Late Cretaceous paleopoles of the two blocks and thus mask or smooth the details of the relative displacement between the Korean Peninsula and the NCB and SCB (Figure 10). In summary, a comparison of the paleomagnetic poles of the CVF with respect to more stable Chinese cratons suggests that neither significant shortening nor rotations are necessary.

As far as the present study is concerned, the paleomagnetic poles of the CVF show excellent coherence with coeval poles from China, suggesting that Korea has been rigidly attached to China (by implication to Eurasia) at least since the Cretaceous (Figure 10). However, earlier paleomagnetic studies on Mesozoic to Tertiary rocks from the Korean Peninsula showed vertical rotations with respect to China (e.g., Lee et al. 1987; Kim and Van der

Voo 1990; Doh and Piper 1994; Zhao et al. 1994, 1999; Doh et al. 1999, 2002; Lee et al. 1999; Zhao et al. 1999; Uno 2000; Park et al. 2003, 2005, 2007; Kim et al. 2009). Then, viable solutions to explain the various degrees of vertical rotations observed in previous paleomagnetic investigations required that individual geologic terranes of the Korean Peninsula formed in Mesozoic to Tertiary experienced different degrees of vertical rotations.

To evaluate the tectonic evolution of the Korean Peninsula, temporal evolution of vertical rotations between the observed and expected magnetic declination for all the available paleopoles documented in the peer-reviewed literature with modern paleomagnetic standards was displayed (Figure 11). To estimate vertical rotation for individual study, three commonly used reference poles from China were assigned for Neogene (Zhao et al. 1994), Paleogene to Cretaceous (Huang et al. 2007), and Jurassic (Gilder et al. 1997). We limited vertical rotation analysis extending only up to Jurassic because the collision

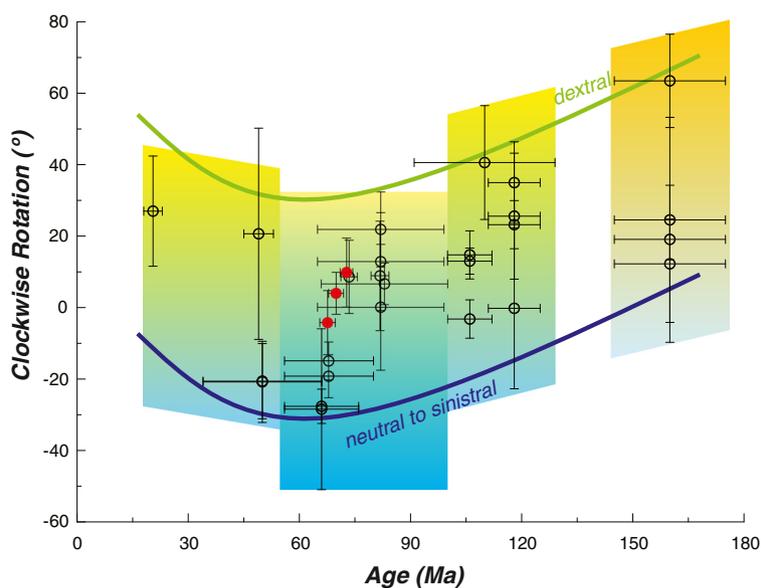


Figure 11 Variation of temporal clockwise rotations of the Korean Peninsula. Vertical rotations were calculated by comparing the observed and expected magnetic declination with respect to the three commonly used reference poles from China for Neogene (Zhao et al. 1994), Paleogene to Cretaceous (Huang et al. 2007), and Jurassic (Gilder and Courtillot 1997). Results from this study (red circles) indicate no significant rotation.

between NCB and SCB was completed by the middle Jurassic as their paleopoles are overlapping each other (Huang and Opdyke 1991; Gilder et al. 1996; Gilder and Courtillot 1997). If the Korean Peninsula was exposed to constant tectonic forcing rather than episodic tectonic processes since the Jurassic, then a linear variation of vertical rotation is expected. Instead of a linear relation, temporal evolution of vertical rotation tracks a half-cycle of sinusoidal variation with the minimum clockwise rotation at 60 to 70 Ma (Figure 11). Prior to 60 to 70 Ma, the amount of clockwise rotations with respect to Chinese reference poles steadily decreased with time (Figure 11). The compiled data further suggest that the mode of clockwise rotations was even reversed to counterclockwise at the end of the late Cretaceous. Then, a new regime of clockwise rotations began to recover (Figure 11).

Plate motions would remain rather constant over long time intervals. However, major tectonic changes, such as reorganization of neighboring plate motions, alteration of tectonic boundary, and opening of an ocean basin, would leave small circle swath with respect to reference Euler poles. Such cusps might be reflected as changes in local vertical rotations (Figure 11). Then, what triggered a sinusoidal swing of vertical rotations in the Korean Peninsula at 60 to 70 Ma? Possible sources of tectonic contributor would be the two major fault systems of the Tancheng-Lujiang fault and the Korea-Taiwan strait fault (several hundred kilometers away), the nearest subduction front of the Pacific Plate (over 1,000 km away), and far-field expression of the India-Asia collision (over 4,000 km away) in order of increasing distance.

The first and nearest contributor may be the major faults developed in East Asia. Indeed, the Korean Peninsula is bounded by two major faults of the Tan-Lu fault to the northwest and the Korea-Taiwan strait fault to the southeast (Figure 1 inset). Both fault systems are northeast-southwest trending sinistral in motion (Grimmer et al. 2002; Wang 2006). Such mega-sized sinistral faults could have facilitated clockwise rotations of the Korean Peninsula with respect to China, as in synthetic R-type Riedel shear. Diminishing clockwise motion since the Jurassic fits well with the period of active motion of the Tan-Lu fault (Klimetz 1983). It should be noted that similar clockwise rotations of paleomagnetic poles were also observed in Benxi area, China (Uchimura et al. 1996).

The second possible contributor would involve changes in near-field tectonic forcing of the subducting Pacific Plate beneath the Eurasian Plate. Although the exact timing and location of Izanagi ridge subduction is far from being perfectly resolved (Whittaker et al. 2007; Seton et al. 2012), the Izanagi-Pacific ridge system is estimated to have existed from 120 to 60 Ma (Rowley 2008; van der Meer et al. 2012). Timing of the disappearance of the Izanagi-Pacific ridge system is consistent with the change in the plate motion of the Pacific from northwest (older than 62 Ma) to west (younger than 62 Ma) (Butterworth et al. 2014). It is feasible that westerly enhanced motion of the Pacific Plate around 60 to 70 Ma restored clockwise rotations of the Korean Peninsula as the westerly subduction would accumulate a dextral shear along the NE trend fault (Figure 11).

Such a clockwise swing was probably weakened as the new back-arc ocean is created between Japan and Korea in Miocene (Figure 11).

A third probable contributor would be far-field expressions of the India-Asia collision (Acton and Gordon 1989; Besse and Courtillot 1991; Vandamme et al. 1991; Leech et al. 2005). As initially proposed by Zhao et al. (1999), the Kula-Eurasia collision may influence the vertical rotations of the Korean Peninsula.

Conclusions

On the basis of paleomagnetic analysis of the CVF, this study demonstrates the following facts:

1. The weighted mean ages obtained by U-Pb zircon dating from the base/top of the CVF are 72.8 ± 1.7 Ma/ 67.7 ± 2.1 Ma, respectively. Hence, the CVF reflects late Cretaceous volcanism that predates the Paleogene.
2. Reliable paleomagnetic information was extracted from the late Cretaceous ignimbrites in Korea. Paleomagnetic mean directions and paleopoles are $D/I = 19.1^\circ/49.2^\circ$ ($\alpha_{95} = 4.2^\circ$, $k = 76.5$) and $VGP = 73.1^\circ\text{N}/232.1^\circ\text{E}$ ($A_{95} = 3.7^\circ$, $N = 3$) for A1, $D/I = 24.9^\circ/52.9^\circ$ ($\alpha_{95} = 5.9^\circ$, $k = 61.7$) and $VGP = 69.4^\circ\text{N}/217.3^\circ\text{E}$ ($A_{95} = 5.6^\circ$, $N = 11$) for A3, and $D/I = 10.9^\circ/50.1^\circ$ ($\alpha_{95} = 5.6^\circ$, $k = 38.6$) and $VGP = 79.8^\circ\text{N}/242.4^\circ\text{E}$ ($A_{95} = 5.0^\circ$, $N = 18$) for A4.
3. VGP positions of the CVF are statistically identical to the reference poles from China in late Cretaceous to early Paleogene, indicating that Korea has been rigidly attached to China (by implication to Eurasia) at least since the Cretaceous.
4. Earlier paleomagnetic investigations on Mesozoic to Tertiary rocks from the Korean Peninsula suggest that the degrees of clockwise rotations weakened since the mid-Jurassic. Such variation might result from the influence of two major nearby strike-slip faults (the Tancheng-Luijiang fault to the northwest and the Korea-Taiwan strait fault to the southeast), near-field tectonic forcing of subducting Pacific Plate, and far-field expressions of the India-Asia collision.

Competing interests

The authors declare that they have no competing interests.

Authors' contributions

DJ carried out the rock magnetic measurements and Ar/Ar dating and drafted the manuscript. YY carried out the paleomagnetic measurements, statistical analysis, and petrographic description and managed the manuscript. SD and DS conceived of the study and participated in its design and coordination. JK was in charge of the U/Pb dating and chronologic interpretation. All authors read and approved the final manuscript.

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