Response of an African tilapia (*Oreochromis hunteri*) to long-term environmental fluctuations: an integrated approach combining the fossil record and modern-day system

Respons van een Afrikaanse tilapia (*Oreochromis hunteri*) op langetermijn omgevingsschommelingen: een geïntegreerde aanpak die het fossiel archief en huidig systeem combineert

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Summary

Studying the processes that underlie organismal diversity is of major importance to understand the origin of new species. Lineages that have speciated rapidly within a geographically confined area, for example in lakes or on oceanic islands, provide valuable model systems to study these processes. Cichlid fishes inhabiting the Great East African Lakes have thus been studied extensively as a prime example of an adaptive radiation, as they have adapted rapidly to a variety of ecological niches and speciated accordingly. However, the lack of a good fossil record has hampered long-term eco-morphological studies in cichlids. In this respect, abundant fossil cichlid teeth, scales and bones preserved in lake sediments may provide a continuous fossil archive for cichlid species inhabiting a lake, but these remain relatively unexplored.

In this PhD thesis, I studied a 25,000-year old fossil cichlid record preserved in the sediments of isolated crater lake Chala (Kenya/Tanzania). This lake has been prone to strong lake-level changes during this period, which presumably impacted habitat availability for its only endemic cichlid, the tilapiine *Oreochromis hunteri*. As oral tooth shape can function as a proxy for past ecology and may thus reflect past changes in habitat exploitation, I hypothesised that these lake-level changes may be visible in the shape of *O*. *hunteri*'s fossil oral teeth. Our aim was thus to trace long-term morphological changes in oral teeth in response to past lake-level fluctuations.

As this study was the first of its kind, I first had to develop a quantitative method that captures the morphology of individual oral teeth. I therefore compared the performance of two geometric-morphometric techniques, elliptic Fourier analysis (EFA) and semi-landmark analysis (SLM), and found that SLM performed slightly better. Combined with multivariate analysis, this method was capable of distinguishing between oral teeth from two congeneric species (*O. hunteri* and *Oreochromis* cf. *korogwe*), and was successfully used to assign a small set of fossil oral teeth to *O. hunteri*.

Moreover, thorough understanding of a fossil record requires adequate knowledge of its modern-day representatives, yet very little was known about the ichthyofauna currently inhabiting Lake Chala. Part of this thesis therefore focussed on resolving phylogenetic relationships and assess feeding ecology of extant Lake Chala cichlids. Although O. hunteri is the only indigenous cichlid species inhabiting Lake Chala, two other tilapiine species, O. cf. korogwe and Coptodon rendalli, were introduced in the second half of the 20^{th} century. Hybridisation with O. cf. koroque may have compromised O. hunteri's genetic and morphological integrity and could undermine its potential as a modern-day reference frame. I therefore analysed the species integrity of O. hunteri using a combination of overall body morphology and mitochondrial DNA (mtDNA) analysis. I found no signs of past or ongoing hybridisation between O. hunteri and the introduced O. cf. koroqwe, and concluded that the morphological integrity of O. hunteri is currently intact. Moreover, using mtDNA from native *Oreochromis* species from nearby Lake Jipe and man-made Nyumba ya Mungu reservoir, Oreochromis jipe was established as the closest relative of O. hunteri.

In order to interpret the ecological significance of past oral tooth shape changes in *O. hunteri*, an assessment of their current use was crucial. Cichlid oral teeth are mostly of importance in obtaining food, hence I analysed the diet of *O. hunteri* using a combination of gut content and stable isotope analysis. I found that this species retains an herbivorous/detritivorous diet throughout the year, but shows seasonal variation in gut content. By comparing gut content composition to monthly fluctuations in the composition of limnetic phytoplankton and to the composition of littoral food source samples, I determined that *O. hunteri* feeds in the littoral zone of Lake Chala for most of the year, but migrates to the limnetic zone in response to an extensive phytoplankton bloom between July and September. I found similar trends in the guts of the introduced *O. cf. korogwe* and *C. rendalli*, which may thus be competing for food with the endemic *O. hunteri*. However, stable isotope composition revealed significant interspecific differences, likely reflecting a discrepancy in the exploitation of the phytoplankton bloom. I thus concluded that *O. hunteri* currently retains a competitive advantage over the introduced newcomers.

As Oreochromis hunteri spends most of the year feeding close to the shore of Lake Chala, I also aimed to establish how individual teeth are eventually buried at the lake's depositional centre, which is approximately one kilometre away from the nearest shoreline. I collected a short sediment core (27 cm) near shore and described fossil tooth distribution throughout this core in relation to sedimentological properties and the distribution of other biological fossils. I found very high abundances of fossil teeth in the large-grained lower sections of this core, indicating that the majority of fossil teeth is permanently buried in the lake periphery. Moreover, the strong correlation with grain size and other dense fossil remains suggests that the hydrostatic properties of fossil teeth do not allow their transport over very long distances, even during occasional extreme events induced by exceptionally intense mixing of the water column. I therefore surmised that teeth retrieved from cores collected at the centre of the lake were deposited *in situ*.

Finally, I analysed both extant and fossil oral tooth shape of Oreochromis hunteri. I quantified oral tooth shape in twelve O. hunteri specimens and observed a gradual shift in oral tooth shape linked to body size. Small- and medium sized specimens had a bicuspid outer series, followed by multiple series of tricuspids, and these were gradually replaced by unicuspid teeth in very large specimens, possibly reflecting changes in feeding behaviour or territoriality. Moreover, I found a strong linear relationship between tooth size and body size, indicating that fossil tooth size can function as a proxy for past body size. I analysed six fossil assemblages from alternating high- and low lake-level stands in the last 25,000 years and found that the morphospace of fossil and extant oral teeth overlapped completely. I thus concluded that O. hunteri is the only species that has inhabited Lake Chala in the last 25,000 years, and that it has retained its overall ecology throughout this period. However, significant shifts in the proportion of unicuspid teeth were observed between different fossil assemblages. Specifically, a 10% increase in the proportion of unicuspid teeth during low lake-level stands was not linked to an increase in median body size. I interpreted this as a shift in the ontogenetic trajectory of oral tooth development in O. hunteri, in which the species obtains 'adult-like tooth shapes earlier in life when lake levels are low, likely in response to a more extensive soft-bottom habitat. I concluded that O. hunteri has been

coping with climate-driven environmental change through rapid morphological responses in oral dentition. Moreover, it has responded repeatedly in the same way to similar stimuli over the past 25,000 years, thus providing evidence of the morphological versatility of *O. hunteri* over extended periods of time.

This study demonstrates the potential of fossil cichlid remains preserved in lake sediments. I successfully used fossils to trace long-term morphological trends in response to severe environmental change in the form of lake-level fluctuations, which have been considered an important driver of cichlid evolution. Studies of continuous fossil records in lakes with more complex species flocks may elucidate long-term patterns of eco-morphological change on much larger scales, and may further enhance our understanding of the processes that underlie cichlid species diversity. This thesis provides a potential approach for these studies.

Samenvatting

Het bestuderen van processen die aan de basis van de huidige biodiversiteit liggen is van belang om te begrijpen hoe nieuwe soorten ontstaan. Goede modelsystemen voor deze studies zijn taxa die zeer snel nieuwe soorten hebben gevormd binnen een geografisch afgebakend gebied, zoals in meren of op oceanische eilanden. De cichliden die de Afrikaanse Grote Meren bevolken zijn om die reden uitgebreid bestudeerd. Deze vissen zijn een typevoorbeeld van een adaptieve radiatie, omdat ze zich zeer snel hebben aangepast aan een waaier van ecologische niches, en al doende nieuwe soorten hebben gevormd. Maar hun gebrekkige fossiele record maakt het moeilijk om eco-morfologische patronen te bestuderen over zeer lange periodes. Fossiele cichlidentanden, -botten en -schubben die werden bewaard in meersedimenten kunnen een ononderbroken archief bieden van de soorten in dat meer, maar werden tot op heden slechts beperkt onderzocht.

In deze doctoraatsthesis werd een 25000 jaar oud fossiel record van een cichlide bestudeerd in de meersedimenten van het kratermeer Lake Chala (Kenia/Tanzania). Het meerniveau van Lake Chala heeft in die periode sterk geschommeld, en die schommelingen hebben hoogst waarschijnlijk een invloed gehad op de beschikbaarheid van habitats voor de enige endemische cichlidensoort in het meer, de tilapia *Oreochromis hunteri*. De vorm van de orale tanden kan iets zeggen over de ecologie van deze soort, en hoe die in het verleden is gewijzigd na veranderingen in beschikbaar habitat. De hypothese was daarom dat meerniveauschommelingen in Lake Chala zichtbaar kunnen zijn in de vorm van fossiele orale tanden van *O. hunteri*. Om die reden werd getracht om langetermijn morfologische veranderingen in fossiele orale tanden op te sporen

die in verband stonden met meerniveauschommelingen in het verleden.

Vergelijkbare studies waren niet beschikbaar, dus er werd eerst een kwantitatieve methode ontwikkeld die de vorm beschrijft van individuele orale tanden. Twee geometrisch-morfometrische technieken werden daartoe vergeleken: elliptische Fourier analyse (EFA) en semi-landmark analyse (SLM). SLM bleek beter in staat om vormvariatie accuraat te beschrijven en werd gecombineerd met multivariate data-analyse. De methode kon een onderscheid maken tussen de orale tanden van twee nauw gerelateerde soorten (*Oreochromis hunteri* en *Oreochromis* cf. *korogwe*), en kon een beperkte dataset van fossiele orale tanden toewijzen aan *O. hunteri*.

Een goed begrip van een fossiel record van een soort steunt op een goede kennis van de huidige populatie van die soort. Bij aanvang van het onderzoek was er echter zeer weinig gekend over de visfauna die op dit moment in Lake Chala voorkomt. Een deel van deze thesis behandelde daarom de fylogenetische relaties van de vissoorten in het meer, en bestudeerde hun huidige voedselvoorkeuren. O. hunteri is de enige vissoort die van nature voorkomt in Lake Chala, maar twee andere tilapia-soorten werden ergens in de tweede helft van de 20^e eeuw ingevoerd: O. cf. koroqwe en Coptodon rendalli. Hybridisatie met O. cf. koroque heeft mogelijks zowel de genetische als de morfologische integriteit van O. hunteri beïnvloed, en zou de bruikbaarheid van deze soort als huidige vertegenwoordiger van het fossiel record in het gedrang kunnen brengen. Daarom werden anlyses van lichaamsvorm en mitochondriaal DNA (mtDNA) gecombineerd om de integriteit van O. hunteri als soort na te gaan. Er werden geen tekenen gevonden van vroegere of huidige hybridisatie tussen O. hunteri en de geïntroduceerde O. cf. korogwe, en de morfologische integriteit van O. hunteri is dus intact. Bovendien werd ook het mtDNA van andere Oreochromis soorten uit de nabijgelegen meren Lake Jipe en Nyumba ya Mungu geanalyseerd, en bleek Oreochromis jipe de dichtst gerelateerde soort aan O. hunteri.

Voor een correcte interpretatie van vroegere veranderingen in de vorm van orale tanden is kennis van hun huidige functie van belang. Over het algemeen worden orale tanden in cichliden gebruikt voor het vergaren van voedsel, dus werd het dieet van *O. hunteri* bestudeerd met een combinatie van tellingen van maaginhouden en stabiele isotopenanalyse. Deze soort blijkt een herbivoor/detritivoor dieet te hebben doorheen het hele jaar, maar vertoont seizoenale variatie in de maaginhoud. Door deze maaginhouden te vergelijken met de samenstelling van limnetisch fytoplankton en van littorale voedselbronnen kon ik vaststellen dat *O. hunteri* het grootste deel van het jaar in de littorale delen van het meer voedsel vergaart, maar naar de limnetische zone migreert wanneer er daar een fytoplanktonbloei plaatsvindt, jaarlijks in de maanden juli, augustus en september. Ik vond gelijkaardige seizoenale trends in de gentroduceerde *O. cf. korogwe* en *C. rendalli*, die dus mogelijks in competitie treden voor voedsel met *O. hunteri*. De stabiele isotopenanalyse leverde echter significante verschillen op tussen de soorten, die waarschijnlijk het resultaat zijn van de verschillende manieren waarop deze soorten gebruik maken van de jaarlijkse fytoplanktonbloei. Op dit moment heeft *O. hunteri* waarschijnlijk een competitief voordeel ten opzichte van de ingevoerde soorten.

Dit betekent echter wel dat O. hunteri zich het merendeel van de tijd dicht bij de kust bevindt in Lake Chala, en dus werd ook getracht om te achterhalen hoe individuele tanden uiteindelijk in het midden van het meer terecht komen, waar de bestudeerde boorkernen werden ingezameld. Een korte boorkern (27 cm) werd daartoe genomen nabij de kust, en er werd een vergelijking gemaakt tussen de distributie van fossiele tandjes en andere biologische fossielen enerzijds, en de textuur van het sediment anderzijds. Er werden zeer hoge abundanties van fossiele tanden aangetroffen in grofkorrelige secties onderaan de boorkern, die erop wijzen dat het merendeel van de tanden die in het meersediment terecht komen permanent worden begraven aan de rand van het meer. Bovendien werd er een sterke positieve correlatie waargenomen tussen korrelgrootte en de abundanties van fossielen met een hoge densiteit, die er op wijst dat de hydrostatische eigenschappen van fossiele tanden geen transport over grote afstanden toelaten. De fossiele tanden die werden aangetroffen in de lange boorkernen die ver uit de kust werden ingezameld werden dus waarschijnlijk ter plaatse afgezet.

Uiteindelijk werd zowel de huidige als de fossiele orale tandvorm van *O. hunteri* in detail bestudeerd. De orale tanden van twaalf *O. hunteri* specimens werden geanalyseerd, en er werd een graduele verschuiving waargenomen in tandvorm die is gelinkt aan de grootte van het individu. Kleinere specimens hebben een bicuspide eerste rij orale tanden, gevolgd door meerdere rijen tricuspide tanden. Die worden echter geleidelijk aan vervangen door unicuspide tanden in zeer grote individuen, wat mogelijks een verschuiving in hun dieet

of in territoriaal gedrag betekent. Er was ook een sterke lineaire correlatie tussen tandgrootte en lichaamsgroote, zodat de grootte van fossiele tandjes kon worden gebruikt als proxy voor de grootte van de vis waaruit ze afkomstig zijn. Vervolgens werden zes fossiele assemblages bestudeerd die afwisselende hoogen laagstanden vertegenwoordigen in de afgelopen 25000 jaar. De vorm van de orale tanden uit deze assemblages overlapte volledig met tandvorm in de huidige O. hunteri, zodat ik kon concluderen dat deze soort de enige aanwezige vissoort is geweest in Lake Chala in deze periode, tenminste vóór de recente introducties. Deze overlap wees er ook op dat de algemene ecologie van O. hunteri in Lake Chala zeer gelijkaardig is gebleven. Er werden echter wel significante verschuivingen aangetroffen in de proportie unicuspide tanden tussen de verschillende assemblages. De proportie unicuspen steeg met ongeveer 10% in laagstand-assemblages, zonder gepaard te gaan met een stijging in gemiddelde lichaamsgrootte. Deze waarneming werd genterpreteerd als een verschuiving in het ontogenetische traject van tandontwikkeling in O. hunteri. De soort kreeg op die manier reeds 'adulte' tandtypes bij een kleinere lichaamsgrootte wanneer het meerniveau laag was, mogelijks als respons op een meer uitgestrekt zanderig habitat. Dit toont aan dat O. hunteri het hoofd heeft geboden aan klimaatgedreven omgevingsveranderingen door een snelle morfologische respons in orale tandvorm. Bovendien heeft de soort zich herhaaldelijk op dezelfde manier aangepast in de afgelopen 25000 jaar, wat de morfologische veranderlijkheid van O. hunteri over zeer lange periodes bevestigt.

Deze studie toont het potentieel aan van fossiele cichlidenresten die worden bewaard in meersedimenten. Fossiele orale tanden werden met succes gebruikt om langetermijn verschuivingen in morfologie op te sporen die het gevolg zijn van klimaatgedreven omgevingsveranderingen. Meerniveauschommelingen worden bovendien beschouwd als een belangrijke drijfveer van evolutie in cichliden in het algemeen, en onze observaties suggereren een potentieel mechanisme dat snelle adaptatie aan een nieuwe omgeving mogelijk maakt. Het bestuderen van gelijkaardige ononderbroken fossiele records in meren met een complexere cichlidengemeenschap kan langetermijn patronen in eco-morfologie blootleggen op veel grotere schaal, en kan op die manier bijdragen aan een beter begrip van de processen aan de basis van de huidige soortendiversiteit. Deze thesis beschrijft hoe zo'n studie kan worden aangepakt.

CHAPTER 1

General introduction

5 1.1 Cichlid adaptive radiation

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The study of diverse yet relatively young species assemblages that have radiated 6 in a geographically confined area has been essential to understand the causes and 7 consequences of speciation (Coyne & Orr, 2004). Especially adaptive radiations, 8 in which the exploitation of and adaptation to a variety of different environments g results in species differentiation (Schluter, 2000; Losos, 2011), have been used 10 to document the origin and proliferation of species. Well-known examples of 11 terrestrial species radiations include *Anolis* lizards in the Caribbean islands 12 (e.g., Williams, 1972; Jackman et al., 1997; Losos, 2011; Mahler et al., 2013), 13 Darwin's finches on the Galápagos (e.g., Lack, 1947; Boag & Grant, 1981; Grant 14 & Grant, 1989; Lamichhaney et al., 2015) and the Hawaiian silversword alliance 15 (e.g., Carlquist et al., 1965; Carlquist, 1980; Carr, 1985; Baldwin & Robichaux, 16 1995). Undoubtedly, the most dramatic case of aquatic species radiation is 17 provided by the cichlid fishes (e.g., Fryer & Iles, 1972; Stiassny & Meyer, 1999). 18 These tropical freshwater fishes occur in Africa, Madagascar, India and southern 19 and central America, and, with an estimated total of 2300 species, they comprise 20 a considerable part of worldwide vertebrate diversity (Van Steenberge, 2014). 21 Especially in the East African lakes Malawi, Tanganyika and Victoria, cichlids 22 have radiated independently into exceptionally diverse species flocks (Fryer 23

& Iles, 1972), which nevertheless share many phenotypes due to convergence 24 (e.g., Kocher et al., 1993; Muschick et al., 2012). The evolutionary mechanisms 25 underlying cichlid diversity are slowly being unveiled by a combination of 26 information that is embedded in species' genetic codes (e.g., Meyer, 1993; 27 Brawand et al., 2014), detailed morphological analysis (e.g., Elmer et al., 2010a; 28 Muschick et al., 2012), ecological assessment (e.g., Fryer, 1959a; Greenwood 20 & Gee, 1969; Reinthal, 1990) and behavioural studies (e.g., Goodwin et al., 30 1998; Verzijden et al., 2010). Based on these studies, a complex interplay of 31 natural selection, hybridisation and sexual selection is now suggested to be 32 responsible for the observed organismal diversity (Salzburger, 2009; Salzburger 33 et al., 2014). 34

Most studies on cichlid fishes have focused on evolutionary mechanisms 35 within historical times, because the lack of a continuous fossil record hampers 36 examination of long-term patterns in morphology and their underlying processes 37 (Murray, 2001). In this respect, cichlid fossils such as teeth, scales and bones 38 preserved in lake sediments may provide a continuous fossil record representing 39 a lake's past ichthyofauna and have been tentatively explored in Lake Malawi 40 (Reinthal et al., 2011) and more recently Lake Victoria (Muschick et al., 2018). 41 Especially fossil teeth could function as a proxy for niche exploitation, as 42 they reflect past feeding habits. The study of cichlid fossils in lake sediments 43 could thus greatly enhance our understanding of cichlid evolution on much 44 longer timescales, but the great complexity of many cichlid communities 45 and complicated geological history of the lakes they inhabit obstruct easy 46 interpretation of these fossil records. Ideally, a small system with limited species 47 diversity and a continuous fossil record is studied first. In this dissertation, 48 I study the potential of such a fossil record in isolated crater lake Chala. 49 Moreover, the well-resolved climatic past of Lake Chala enables tracing the 50 morphological response of cichlid fossils to past environmental change, in 51 particular strong lake-level fluctuations. 52

53 1.2 Ecological opportunity

Ecological opportunity provided by the availability of complex and discordant
habitats is one component that has been considered crucial for the diversification
of cichlid species (Huber *et al.*, 1997; Danley & Kocher, 2001). The dynamic

geological history and paleoclimate of the East African region has greatly 57 influenced the availability of such habitats in the past (e.g., Rossiter, 1995; 58 Danley et al., 2012). During the Quaternary period, this region has witnessed 59 alternating wet and dry cycles linked to fluctuations in monsoon rainfall due to 60 variations in solar insolation forcing (e.g., Verschuren et al., 2009; Lyons et al., 61 2015), with obvious effects on African palaeohydrology and lake levels (De 62 Cort et al., 2017). In the African Great Lakes (Fig. 1.1), lake-level fluctuations 63 caused extreme contraction (Lake Malawi; Delvaux, 1995; Scholz et al., 2007; 64 Lyons et al., 2011), fragmentation (Lake Tanganyika; Scholz & Rosendahl, 1988; 65 Rossiter, 1995) or even desiccation (Lake Victoria, Johnson et al., 1996), and 66 have had a profound impact on the lakes' biota (Cohen et al., 2007; Danley et al., 67 2012). Indeed, rapid population expansion of inshore demersal haplochromine 68 cichlids in Lake Malawi has been linked to the lake rising to current levels 69 (Genner & Turner, 2014), and fragmentation of Lake Tanganyika is reflected 70 in the phylogeographic patterns of its rock-dwelling cichlids (Sturmbauer & 71 Meyer, 1992; Verheyen et al., 1996; Rüber et al., 1999; Nevado et al., 2013). 72 The relatively recent desiccation of Lake Victoria ($\sim 14,000$ yr BP; Johnson 73 et al., 1996), which is nevertheless populated with an extremely diverse cichlid 74 species flock (\sim 500 species; Snoeks, 2000), is perhaps the most striking example 75 of the ecological opportunity created by these lake-level changes. 76

At least in the case of Lake Tanganyika, lake-level fluctuations have been 77 considered to have functioned as a 'species pump'. Fragmentation of this 78 lake into several sub-basins due to strong lake-level drops facilitated allopatric 79 speciation among cichlids, greatly enhancing species diversity even when lake 80 levels rose again and reunited communities from different sub-basins (Rossiter, 81 1995). More complex effects of lake-level fluctuations on speciation have been 82 suggested for Lake Malawi. Based on phylogenetic patterns in Malawi cichlids, 83 Ivory et al. (2016) hypothesised that an expansion of rocky habitats allowed 84 their diversification during past periods of high lake level. In addition, lake 85 lowstands were accompanied by more turbid conditions, leading to episodes 86 of extinction but also to interspecific hybridisation, which is known to be an 87 important source of evolutionary novelty for selection to act upon (Verheyen 88 et al., 2003; Seehausen, 2004; Meier et al., 2017a). Thus, both lake-level states 89 are suggested to have contributed to the phenomenal species diversity observed 90 in African cichlid fishes today. 91

1

⁹² 1.3 The trophic apparatus as key innovation

Ecological opportunity alone does however not predict whether cichlid lineages 93 radiate or not. Several lineage-specific traits have been proposed to be of 94 importance for diversification and speciation in cichlid fishes. The evolutionary 95 potential of cichlids is likely contingent on unique genomic properties (See-96 hausen, 2006; Brawand et al., 2014), and the evolution of egg spots (Hert, 97 1989; Salzburger et al., 2005), maternal mouthbrooding (Greenwood, 1979; 98 Sefc, 2011) and a sophisticated trophic apparatus (Liem, 1973; Hulsey et al., 90 2006) have all been proposed to have promoted their rapid speciation. The 100 trophic apparatus has even been advanced as a 'key innovation' (Liem, 1973; 101 Hulsey et al., 2006), as it provides a clear tool for ecological adaptation and 102 niche diversification, eventually contributing to adaptive speciation (Schluter, 103 2000).104

The evolution of jaws has been a major driver of vertebrate evolution in general (Mallatt, 1996; Liem *et al.*, 2001), and the diversification of trophic morphology in cichlids is a great example of the evolutionary potential of this trait (Fryer & Iles, 1972). Cichlid fishes are not unique in their possession of a

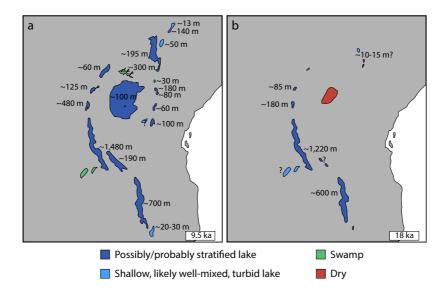


Figure 1.1: Palaeogeographic maps of East Africa showing lake-level stands 9,500 (a) and 18,000 (b) year BP. Numbers represent estimated water depths. Modified from Salzburger *et al.* (2014)

second set of jaws (i.e. pharyngeal jaws; Schaeffer & Rosen, 1961; Liem, 1986; 109 Wainwright, 1989), but the fusion of the pharyngeal plates and the modified 110 musculature are unique innovations that have enhanced this jaw's mobility and 111 led to a functional decoupling from the oral jaws (Liem, 1973; Meyer et al., 112 1990; Hulsey et al., 2006; Fig. 1.2a). This way, the oral jaws were released to 113 specialise on prey collection while the pharyngeal jaws do the prey processing 114 (Liem, 1973; Lauder et al., 1989). The resulting independent divergence of 115 jaws, combined with diverging tooth shapes, has produced remarkable levels 116 of morphological specialisation (Fryer & Iles, 1972; Hulsey et al., 2006) and 117 facilitated extraordinary cases of fine-scaled resource partitioning (Fryer, 1959a; 118 Reinthal, 1990). Nevertheless, convergence of trophic structures is common in 119 cichlids, both within (Rüber et al., 1999; Rüber & Adams, 2001) and between 120 water bodies (Kocher et al., 1993; Muschick et al., 2012), even across continents 121 (Winemiller et al., 1995). That similar ecological problems generate similar 122 ecomorphological solutions indicates that certain morphological constraints 123 may influence the direction of evolution (Albertson & Kocher, 2006), and 124 suggests that natural selection is shaping phenotypic diversity. 125

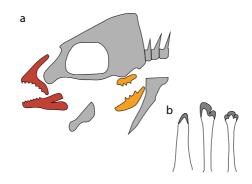


Figure 1.2: a Schematic representation of oral (red) and pharyngeal (yellow) jaws in cichlid fishes, modified from Machubi *et al.* (2007). b Schematic drawing of oral jaw teeth of *Oreochromis hunteri*, representing the three general tooth types present in cichlid fishes. From left to right: unicuspid, bicuspid and tricuspid tooth

This unique combination of trophic divergence within clades and morphological convergence between clades has therefore stimulated investigations of the functional morphology of cichlid jaws. In this context, the development of a quantitative method to analyse shape (i.e. geometric morphometrics; Bookstein, 1991; Zelditch *et al.*, 2004) has proven extremely valuable in cich1

lid research (reviewed in Kerschbaumer & Sturmbauer, 2011). Such studies 131 have highlighted for example the importance of physical constraints on jaw 132 morphology (Hulsey, 2006), the presence of structural trade-offs (Tkint et al., 133 2012), and the existence of intraspecific polymorphisms in the lower pharyngeal 134 element (e.g., Trapani, 2004; Muschick et al., 2011). Although the resulting 135 variation in jaw shape is astounding, the diversity in cichlid tooth shapes is 136 equally impressive (Greenwood & Gee, 1969; Fryer & Iles, 1972). Tooth shape 137 variation can be summarised in three major tooth types, based on the number 138 of cusps they bear: unicuspid, bicuspid and tricuspid teeth (Fig. 1.2b). Teeth 139 are often implanted in the jaws in multiple series, in which each separate series 140 bears similarly-shaped teeth (homodonty; Streelman et al., 2003), but shape 141 may differ substantially between series. Subtle differences in cusp shape aid 142 species to specialise in a certain niche, and these have enabled the fine-scaled 143 resource partitioning observed in the Great Lakes' cichlid communities (Fryer, 144 1959a; Greenwood & Gee, 1969; Reinthal, 1990). 145

The developmental pathways and genetic architecture controlling cichlid 146 jaw and tooth morphology are now relatively well-understood (e.g., Huysseune 147 & Sire, 1992; Streelman et al., 2003; Albertson & Kocher, 2006; Hulsey et al., 148 2017; Singh et al., 2017). Increasing numbers of genes underlying the observed 149 craniofacial diversity are being discovered (Albertson & Kocher, 2006; Gunter 150 et al., 2013) and substantial genetic covariance exists between the oral and 151 pharyngeal jaws (Hulsey et al., 2017). Cichlids typically replace their teeth 152 continuously via *de novo* formation of tooth germs (polyphyodonty), and tooth 153 shape can vary considerably throughout a fish's life due to temporal changes in 154 gene expression (Streelman et al., 2003). Such ontogenetic changes are possibly 155 associated with alterations in feeding habits when juveniles reach adulthood, 156 or with sexual maturation. Differences in the timing and rate of ontogenetic 157 trajectories are an important source of phenotypic diversity, and may provide 158 cichlids with the means to rapidly adapt to a changing environment (Meyer, 159 1987; Streelman et al., 2003), as has for example been documented in the visual 160 system of neotropical cichlids (Härer et al., 2017). 161

¹⁶² 1.4 Small systems and non-radiating lineages

Many evolutionary mechanisms and processes have been elucidated by studying 163 the species flocks that populate the Great Lakes of Africa, and their extraordi-164 nary outcome is also best observed in these lakes. Yet the enormous habitat 165 diversity and difficulties in determining the relationships between species have 166 complicated untangling certain principles (e.g., Verheyen et al., 2003; Joyce 167 et al., 2011; Ford et al., 2015; Kavembe et al., 2016; Meyer et al., 2017), and 168 have compelled scientists to investigate the inhabitants of smaller lake systems. 169 Such systems have thus been used to characterise the genomic properties of 170 the early stages of adaptive radiation (Ford *et al.*, 2015; Malinsky *et al.*, 2015; 171 Kavembe et al., 2016; Ronco & Salzburger, 2016), but they have also functioned 172 as 'natural laboratories' to study adaptive responses to, for example, changes in 173 light environments (Härer et al., 2017; Torres-Dowdall et al., 2017). Famously, 174 the most compelling evidence for the enigmatic process of sympatric speciation – 175 divergence in the presence of gene flow – has been found in small, monophyletic 176 species flocks inhabiting isolated crater lakes (Schliewen et al., 1994; Barluenga 177 et al., 2006; Elmer et al., 2010b), although recently new debate has emerged 178 about some of the interpretations made (Martin *et al.*, 2015; Foote, 2018). 179

Small systems harbouring non-radiating cichlid lineages have also been 180 pivotal in explaining which factors influence diversification. Through a com-181 parative analysis of radiating and non-radiating cichlid lineages in both large 182 and small lakes across tropical Africa, Wagner *et al.* (2012) found that a 183 combination of lake depth, solar radiation and sexual dichromatism predicted 184 whether cichlids radiated or not. Not all cichlid lineages thus have the same 185 propensity for radiation. One group of cichlids whose members rarely radiate 186 is the paraphyletic species group formerly referred to as 'Tilapia' (Ribbink, 187 1990; Nagl et al., 2001; Klett & Meyer, 2002; Dunz & Schliewen, 2013), al-188 though small incipient radiations of tilapiine cichlids have been found in some 189 systems (Schliewen & Klee, 2004), often driven by extreme environmental 190 conditions (Ford et al., 2015; Kavembe et al., 2016). Despite sharing many 191 characteristics with radiating cichlid lineages, tilapiines differ fundamentally 192 in certain inherent traits, for example in their genomic properties (Brawand 193 et al., 2014). Moreover, whereas radiations often produce highly stenotopic 194 communities in which each species is perfectly adapted to its relatively narrow 195

ecological niche, tilapia species display exceptionally high levels of versatility
in behaviour, ecology and morphology (Fryer & Iles, 1972; Trewavas, 1983;
Ribbink, 1990).

Tilapiines are indeed remarkably tolerant to wide ranges of temperature, 199 dissolved oxygen, salinity, alkalinity and pH (reviewed in Philippart & Ruwet, 200 1982), and several species occupy a range of different habitats during their 201 lifespan (Trewavas, 1983; Ribbink, 1990). Their diet is generally herbivorous 202 or detritivorous, but changes in feeding habits are commonly reported and 203 have been linked to ageing (Bruton & Boltt, 1975; Whyte, 1975; Gophen, 1980; 204 Njiru et al., 2004), to seasonal differences in food availability (Spataru, 1976; 205 Man & Hodgkiss, 1977; Spataru, 1978a), to breeding status (Spataru, 1978b) 206 or simply to opportunity (Bruton & Boltt, 1975). Moreover, there is a large 207 amount of intraspecific variability in growth and onset of sexual maturation 208 between populations from different localities (Brummett, 1995; Duponchelle 209 & Panfili, 1998; Bwanika et al., 2007), strongly depending on environmental 210 variables. 211

As a result, requirements for tilapia aquaculture are not very demanding, 212 and a considerable part of the global freshwater fish production is currently 213 tilapia-based (FAO, 2014). Especially large Oreochromis species are used 214 for this purpose, such as O. niloticus, O. mossambicus, O. aureus, and their 215 different hybrids. Although part of the fish production is confined to artificial 216 fish ponds, natural systems worldwide have been stocked with alien Oreochromis 217 species to boost fisheries (Eknath & Hulata, 2009), where their versatile nature 218 has allowed them to readily adapt to new circumstances. In many places, 219 these anthropogenic introductions have resulted in ecological disasters. The 220 effects of invasive *Oreochromis* species are typically most apparent in the 221 native ichthyofauna (reviewed in Canonico et al., 2005), but in Mexico for 222 example the introduction of O. niloticus has also been linked to population 223 decline in the iconic yet critically endangered axolotl (Zambrano et al., 2010; 224 Alcaraz et al., 2015). The negative impacts of invasive aliens on a natural 225 system are often the result of a complex interplay of factors (Gozlan et al., 226 2010), and in Oreochromis include direct predation (De Moor et al., 1986) 227 and competition for food (Ogutu-Ohwayo, 1990; Daga et al., 2016) or prime 228 breeding grounds (Canonico *et al.*, 2005). Additional concern is raised by the 229 propensity of introduced *Oreochromis* species to hybridise with indigenous 230

tilapiine species (Agnèse et al., 1998; D'Amato et al., 2007; Nyingi & Agnèse, 231 2007; Angienda et al., 2011; Deines et al., 2014; Ndiwa et al., 2014). The 232 introgression of alien genes into local species may lead to homogenisation of 233 the gene pool and thus induce a loss of genetic diversity (Nyingi & Agnèse, 234 2007; Crispo et al., 2011; Firmat et al., 2013), but it also potentially impacts 235 important morphological features (Parnell et al., 2012; Holzman & Hulsey, 236 2017). Interspecific hybridisation is considered as an important mechanism 237 in generating natural cichlid diversity (Salzburger et al., 2002; Seehausen, 238 2004; Meier et al., 2017a), but from a conservational point of view it creates a 239 worrying prospect (Pullin, 2002). 240

²⁴¹ 1.5 The fossil record of cichlids

Molecular phylogenetic techniques have proven invaluable in elucidating many 242 of the evolutionary processes driving cichlid speciation (e.g., Wagner *et al.*, 243 2012; Brawand et al., 2014), but also in reconstructing the origin of certain 244 biogeographic patterns (e.g., Joyce et al., 2005; Meier et al., 2017a). Moreover, 245 they can be used to estimate divergence times between species based on the 246 principle of the 'molecular clock', which assumes a correlation between the age 247 of species divergence and the number of genetic differences (Zuckerkandl & 248 Pauling, 1962). But large differences exist in the rate of genetic change between 249 lineages (Drummond et al., 2006), and a good calibration of the molecular clock 250 is necessary to accurately estimate the timing of past divergence events. For an 251 estimation of absolute minimum age, calibration is commonly done using fossils, 252 and clocks calibrated in this manner have for example been used to estimate 253 the divergence times of large cichlid species flocks (Genner et al., 2007), or 254 to link phylogeographic patterns to past lake-level fluctuations (Genner & 255 Turner, 2014). However, the cichlid fossil record has thus far proven inadequate 256 to answer such important questions as the timing of divergence between 257 African and neotropical cichlids, so that the long-standing debate between 258 their Gondwanan vicariance or transatlantic dispersal remains unresolved (e.g., 259 Genner et al., 2007; Azuma et al., 2008; Smith et al., 2008; Friedman et al., 260 2013; Matschiner et al., 2017). 261

Proper fossil records are useful in many other ways, notably in extrapolating evolutionary and ecological observations over longer time scales. In

this respect, the study of morphology and wear of fossil teeth has proven 264 effective in diverse vertebrate taxa, such as hominids (reviewed in Ungar, 2017), 265 bovids (e.g., Solounias & Moelleken, 1992; Fortelius & Solounias, 2000; Blondel 266 et al., 2010) and fish (e.g., Purnell et al., 2007; Bellwood et al., 2014; Hundt 267 & Simons, 2018). But for cichlids, the scarcity of continuous fossil deposits 268 and the disarticulated nature of remains has hampered a detailed study of 269 long-term trends in their ecology and evolution (Murray, 2001). The cichlid 270 teeth and scales preserved in lake sediments hold the best hope for a continuous 271 record, and are hence beginning to be explored in Africa's Great Lakes, e.g. in 272 Lake Malawi (Reinthal et al., 2011) and Lake Victoria (Muschick et al., 2018). 273 However, the species richness in these large lakes makes it difficult to assign 274 fossils to specific lineages. Also, the past lake-level fluctuations that presum-275 ably influenced cichlid speciation (cf. above) potentially impacted important 276 taphonomic processes and the depositional environment in which the fish are 277 preserved, impeding straightforward interpretation of these records. In this 278 context, studying the fossil record of cichlid fishes in a small and simple lake 279 system prone to similar environmental dynamics becomes appealing to assess 280 the true potential of this approach. In this thesis, I present the results of such 281 a study on the sediment archive of crater Lake Chala. 282

²⁸³ 1.6 Lake Chala

²⁸⁴ 1.6.1 Local climate & sediment archive

Lake Chala $(03^{\circ}19'S, 37^{\circ}42'E)$ is a 4.2 km² crater lake at the border between 285 Kenya and Tanzania, located on the lower east slope of Mount Kilimanjaro (Fig. 286 1.3a). It is contained within a steep-sided caldera, and the narrow shoreline 287 quickly drops down to a depth of nearly 55 m (Opitz, 2008), from where soft 288 sediments slope more gently towards a maximum depth of around 92 m in 289 the middle of the lake (Fig. 1.3b; Moernaut et al., 2010). The local climate 290 is tropical semi-arid, with mean temperatures varying between 26 and 30°C. 291 Lake-surface evaporation supersedes total annual rainfall (~ 565 mm), but 292 sub-surface in- and outflow strongly influence the lake's water budget and keep 293 it fresh (Moernaut et al., 2010). The latitudinal zone of peak insolation and the 294 intertropical convergence zone (ITCZ) pass across this equatorial region twice 295 per year, resulting in a characteristic bimodal pattern of rainfall seasonality: 296

southeasterly monsoon winds bring 'long rains' from March to May, whereas 297 the northeasterly monsoon winds bring 'short rains' from late October to 298 December. During most of the year, wind-driven turbulence and night-time 299 convection in this sheltered crater basin are relatively modest, limiting frequent 300 oxygenation to the upper 15-30 m of the water column. The surface water 301 is cooled by a combination of strong southerly winds and reduced seasonal 302 insolation during the southern hemisphere winter from late June to September. 303 which promotes convection and expands the mixed layer to between 35 m and 304 45 m (Buckles et al., 2014). This deep mixing brings up dissolved nutrients that 305 have been accumulating in the lower water column and induces an extensive 306 phytoplankton bloom at the end of the long dry season (Barker *et al.*, 2011; 307 Wolff et al., 2011; Buckles et al., 2014). 308

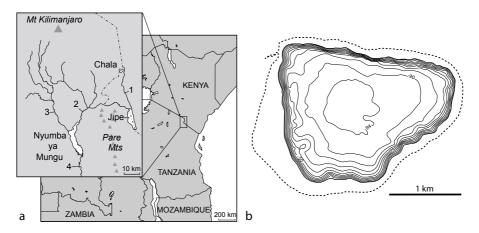


Figure 1.3: Skeleton maps of East Africa and the Upper Pangani River basin (inset) in northern Tanzania, with indication of the sampled surface waters Chala, Jipe and Nyumba ya Mungu (a). River systems are indicated by numbers. 1: Lumi, 2: Ruvu, 3: Kikuletwa, 4: Pangani. Bathymetry of Lake Chala (b) with dotted line representing its catchment. Modified from Moernaut *et al.* (2010)

The sediment record of Lake Chala demonstrates long-term continuity of lacustrine conditions (Verschuren *et al.*, 2009; Moernaut *et al.*, 2010), and is currently the subject of a multifaceted research programme aiming to reconstruct the long-term climate history and ecosystem dynamics of equatorial East Africa (Verschuren *et al.*, 2013). Notably, seismic-reflection data revealed that the lake has experienced pronounced lake-level fluctuations since its origin approximately 250,000 years ago. These fluctuations probably ranged from 1

 \sim 30-40 m below to \sim 10 m above current lake levels during low- and highstands respectively (Moernaut *et al.*, 2010). This change in total lake depth likely

respectively (Moernaut *et al.*, 2010). This change in total lake depth likely had an impact on the availability and distribution of different types of aquatic habitat and food resources, and can thus be expected to have influenced the ecology and evolution of indigenous ichthyofauna in Lake Chala.

321 1.6.2 Oreochromis hunteri, the mountain Chromis

Lake Chala currently harbours the only known population of the Chala Tilapia, 322 Oreochromis hunteri Günther, 1889 (Fig. 1.4a). This was the only fish species 323 present in the lake when A. Günther described it in 1889 as the type species for 324 a new genus: Oreochromis, 'the mountain Chromis', in reference to its home at 325 the foot of Mount Kilimanjaro (Dadzie et al., 1988). Primarily because of its 326 restricted distribution in a single small lake, this endemic species has been listed 327 as 'critically endangered' on the IUCN red list (Bayona et al., 2006), yet very 328 little is known about its ecology. Based on morphological similarities Trewavas 320 (1983) proposed a close relationship to Oreochromis jipe from nearby Lake Jipe, 330 but this hypothesis has never been tested phylogenetically. Similarly, Trewavas 331 (1983) could not elaborate on its feeding habits and breeding behaviour, except 332 that Lowe (1955) reported having observed the young fish feeding on algae and 333 debris from the bottom between the rocks. Indeed, in the modern-day lake 334 system the steep rocky crater walls and open-water (limnetic) zone provide 335 the only oxygenated habitats available to O. hunteri. The rocks are covered 336 by epilithic algae and interspersed with small sandy patches trapping plant 337 litter and other organic detritus, thus providing a dependable food source 338 throughout the year. In contrast, primary production in the limit zone is 339 very low (secchi-disk transparency ranging between 6 and 9 m; van Bree et al., 340 unpublished data) except during the phytoplankton bloom between July and 341 September. However, a detailed quantitative assessment of O. hunteri's diet is 342 still lacking. 343

344

Investigations into the origin and ecology of *Oreochromis hunteri* are especially relevant because it now shares Lake Chala with three anthropogenically introduced cichlid species. These introductions probably took place sometime in the second half of the 20th century (Caxton Oluseno, pers. comm.), most likely after 1980 as surveys carried out on Lake Chala in 1889, 1902, 1946, 1952

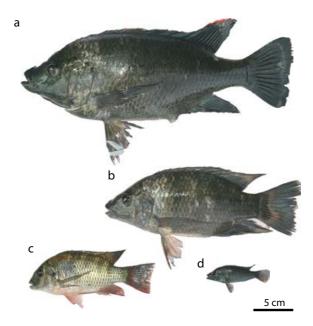


Figure 1.4: The extant ichthyofauna of Lake Chala consists of the endemic Oreochromis hunteri (a), introduced tilapiines Oreochromis cf. korogwe (b) and Coptodon rendalli (c), and the small haplochromine Haplochromis sp. 'Chala' (d)

and 1980 only found O. hunteri present (Günther, 1889; Dadzie et al., 1988). 350 Two tilapiine species have been found in catches since then, being *Oreochromis* 351 koroque Lowe, 1955 (Fig. 1.4b) and Coptodon rendalli Boulenger, 1896 (Fig. 352 1.4c; Dadzie et al., 1988; Seegers et al., 2003), although the former of the two 353 was identified with some doubt. Because of this uncertainty, this taxon is 354 identified as O. cf. koroque in this thesis (see Chapter 3). A small cichlid 355 species identified as *Haplochromis* sp. "Chala" was also introduced in that 356 period (Seegers et al., 2003). These species may interfere with the endemic O. 357 hunteri in multiple ways. First, the congeneric nature of O. cf. korogwe and O. 358 hunteri increases the risk of hybridisation and loss of genetic diversity (Nyingi 359 & Agnèse, 2007; Crispo et al., 2011; Firmat et al., 2013). Second, similar 360 habitat requirements of the introduced tilapiines may lead to the indigenous 361 species being outcompeted, as has happened for example with the indigenous 362 oreochromines of Lake Victoria after the introduction of O. niloticus (Twongo, 363 1995; Goudswaard et al., 2002). 364

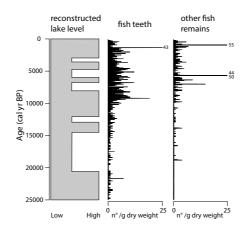


Figure 1.5: Counts of cichlid fossils with respect to reconstructed lake level (after Moernaut *et al.*, 2010) in a sediment core of Lake Chala spanning the last 25,000 years

365 1.6.3 The fossil record of Oreochromis hunteri

Fish remains in the form of fossil teeth, scales and bones are abundantly present 366 in the sediment archive of Lake Chala throughout at least the last 25,000 years 367 (Fig. 1.5). These most probably belong to O. hunteri, or to another species of 368 cichlid that has since gone extinct; the introduced species are not represented in 369 the fossil record except possibly in the uppermost centimetres of the sediment 370 column. However, the seismic-reflection data reveal that considerable lake-level 371 fluctuations have occurred also during the last 25,000 years (Fig. 1.5), with 372 prolonged lowstands primarily during the early late-Glacial period ($\sim 19,500$ -373 14,500 years ago), the Younger Dryas ($\sim 13,000-11,500$ years ago) and much 374 of the mid-Holocene (~8500-3000 years ago). These fluctuations potentially 375 influenced the ecology of O. hunteri by changing the availability of different 376 aquatic niches. Specifically, the lowstands were probably severe enough (\sim 30-40 377 m, Moernaut et al., 2010) to oxygenate soft-bottom areas in the lake periphery 378 at the foot of the rocks, which currently experience permanent anoxia due 379 to their great depth (at least 55 m; Opitz, 2008). Hence, a soft-bottom 380 benthic habitat and associated food source may have become available during 381 such lowstands (Fig. 1.6). Presumably, such a shift would be reflected in 382 ecologically important morphological traits, for example tooth shape. The 383 excellent preservation of fish teeth in the sediment record of Lake Chala creates 384

a unique opportunity to trace such morphological shifts over extended periods
of time, as a proxy for the fishes' changes in niche exploitation in response to
climate-driven lake-level fluctuations.

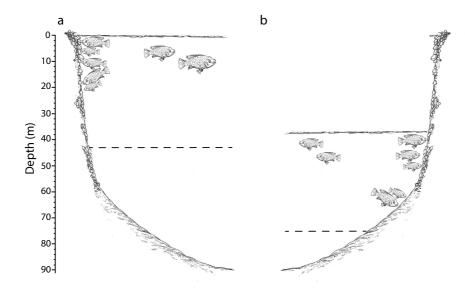


Figure 1.6: Schematic representation of the hypothesised effect of strong lake-level fluctuations on the available habitat for *O. hunteri*. During high lake-level stands, oxygenated habitats are limited to rocks and small sandy patches in the littoral zone, and the open-water (limnetic) zone (a). During lake lowstands, oxygenation of soft sediments near the lake periphery may bring a soft-bottom benthic habitat and associated food sources within reach (b). Dashed line represents the maximum depth of the oxycline during the mixing season. Drawing by Tim Goethals

388 1.7 Outline of this thesis

The principal aim of this PhD thesis is a detailed study of the fossil cichlid teeth 389 preserved in the sediment record of Lake Chala. The isolated nature of the 390 lake, its well-studied environmental past, and the known timing of recent fish 391 introductions provide a uniquely controlled setting for morphological analysis 392 of its fossil record. However, correct interpretation of this fossil record requires 393 thorough understanding of the ecology and genetic relationships of modern-day 394 Oreochromis hunteri, and of the taphonomic processes that might influence the 395 fossil record. Consequently, the research programme of this PhD project has 396

also given due attention to these aspects. The results of this study can thus bestructured in six chapters, as outlined below.

³⁹⁹ 1.7.1 Method development for geometric-morphometrics of fos ⁴⁰⁰ sil fish teeth

Chapter 2 of this thesis describes the development of several methods used 401 in subsequent chapters. Important morphological traits are examined using 402 geometric-morphometrics. Based on existing literature, a set of landmarks 403 is defined to analyse the general body shape of the fishes. These landmarks 404 are then applied to personal collections of the two *Oreochromis* species now 405 inhabiting Lake Chala (O. hunteri and O. cf. koroque) to test the poten-406 tial of this method to discriminate between species. A quantitative method 407 capturing cichlid tooth shape had been developed earlier by Wautier et al. 408 (2002), but required undamaged teeth, which is not always possible in fossil 409 teeth. I therefore develop a method combining length-width measurements 410 of complete teeth with quantitative shape analysis of the tooth top alone (i.e. 411 the enameloid-covered crown). I further compare the outcome of two alterna-412 tive geometric-morphometric techniques in describing the shape of oral tooth 413 crowns: elliptic Fourier analysis (EFA) and semi-landmark analysis (SLM). 414 The optimal technique is then used to compare and describe the tooth shapes 415 of O. hunteri and O. cf. korogwe, and a first set of fossil teeth is added to the 416 dataset to assess the feasibility of this approach. 417

418 1.7.2 Genetics and ecology of the extant ichthyofauna of Lake 419 Chala

Morphological integrity of the trophic apparatus of modern-day O. hunteri 420 is fundamental to its function as modern-day reference for the fossil record. 421 However, hybridisation with recent newcomers may compromise this integrity 422 (Parnell et al., 2012; Holzman & Hulsey, 2017). In Chapter 3, I therefore 423 search for traces of interbreeding by combining whole-body geometric mor-424 phometry with molecular phylogenetic analysis of mitochondrial loci from O. 425 hunteri and O. cf. korogwe. Comparing the obtained sequences to relevant 426 GenBank entries, I aim to establish the true identity of the O. cf. koroque 427 population in Lake Chala. Using fishes from Lake Jipe and Nyumba ya Mungu 428

reservoir, I expand the analysis to all four *Oreochromis* species currently inhab-429 iting the Upper Pangani River system to determine the closest living relative 430 of O. hunteri, and hence the possible source population of the ancestral species 431 which colonised Lake Chala. Based on the observed haplotype diversity, I also 432 intend to provide an estimated timing of arrival for *O. hunteri* in Lake Chala. 433 Oral tooth shape is strongly linked to the feeding habits of a species, yet 434 very little is known about the diet of O. hunteri. The feeding ecology of O. 435 hunteri, O. cf. korogwe and C. rendalli in Lake Chala is therefore studied in 436 **Chapter 4** using a combined analysis of gut content composition and stable 437 isotopes. Monthly collecting of fish over the course of 20 months enables tracing 438 of seasonal variation in gut content, which is expected to occur in relation to 439 the annual phytoplankton bloom. I combine these observations with monthly 440 counts of the limnetic phytoplankton, and with the composition of potential 441 food sources collected from rocks and sandy patches at the lake periphery. 442 This way, I aim to assess the food sources these species are exploiting at 443 different times of the year. The isotopic composition of fish tissue is then 444 compared between species to trace any general interspecific differences in 445 diet, and interpreted in conjunction with observations on gut content. This 446 chapter concludes with a discussion of the consequences of anthropogenic fish 447 introductions on the ecology and conservation status of O. hunteri. 448

449 1.7.3 Sedimentation processes influencing fossil distributions

The results of Chapter 4 suggest that O. hunteri spends most of its life close 450 to shore. Correct interpretation of the fossil record therefore requires a proper 451 understanding of how O. hunteri's teeth end up at the depositional centre of 452 the lake, where the 25,000-year sediment sequence has been extracted. The 453 distribution of fossil teeth throughout a short core collected near the lake 454 periphery is analysed in **Chapter 5**. I link the distribution and size variation 455 of the fish teeth to important sedimentological variables such as the grain size 456 of the clastic-mineral fraction, as well as to the distribution of other types of 457 biological fossils. This may aid in understanding processes of tooth burial and 458 transport, and how these may influence our observations from sediment cores 459 collected offshore. 460

461 1.7.4 Long-term trends in oral tooth morphology of *Oreochromis* 462 *hunteri*

In Chapter 6, I assess the effects of past lake-level fluctuations (Verschuren 463 et al., 2009; Moernaut et al., 2010) on the local O. hunteri population through 464 analysis of fossil tooth shape. The Chapter starts with implementation of 465 the methods developed in Chapter 2 to achieve a thorough description of 466 oral tooth morphology of extant O. hunteri. I then collected oral teeth from 467 six time windows in the last 25,000 years representing alternating high and 468 low lake-level stands. These six fossil assemblages are interpreted using the 469 generated modern-day framework. I explore whether O. hunteri has indeed 470 been the only species present in Lake Chala in this period, and compare 471 morphospace occupation of fossil lowstand and highstand assemblages. This 472 way, I aim to track long-term changes in the oral tooth morphology in relation 473 to climate-driven environmental change. 474

475 **1.7.5** General discussion

In final **Chapter 7**, the findings of all previous chapters (Chapter 2-6) are 476 integrated into a general discussion. I make an assessment of the methods 477 used and discuss their potential application in other systems. Genetics of 478 the oreochromines of Lake Chala are briefly discussed, and tilapiine oral 479 tooth shape is linked to feeding habits. Taphonomic processes influencing 480 offshore fossil tooth distribution in Lake Chala sediments are discussed with 481 respect to observations in the fossil record, and I elaborate on the origin of O. 482 hunteri in Lake Chala and its subsequent eco-morphological response to past 483 environmental fluctuations. 484

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CHAPTER 2

Exploring the methods

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510 2.1 Abstract

The sedimentary archive of Crater Lake Chala in East Africa contains abundant 511 fossil teeth of cichlid fishes throughout at least the last 25,000 years. Here, 512 we use morphometric analyses of oral teeth from the two extant Oreochromis 513 species inhabiting Lake Chala to explore the feasibility of tracing adaptive 514 modification of the cichlid trophic apparatus in the lake's fossil record. We 515 compared the performance of semi-landmark analysis (SLM) and elliptic Fourier 516 analysis in capturing morphological variation in oral tooth crowns, and found 517 that SLM, supplemented with tooth neck measurements, ensured consistency 518 with whole-body landmark analyses. Whole-body and tooth morphology data 519 both allowed to discern between the indigenous *Oreochromis hunteri* and the 520 recently introduced *Oreochromis* cf. koroqwe. Both species have an oral jaw 521 arrangement of outer-series bicuspid and inner-series tricuspid teeth, but O. 522 *hunteri* has generally more slender teeth with a proportionally taller major cusp 523 than O. cf.koroqwe, suggesting fine-scaled resource partitioning. Exploratory 524 analysis of three fossil samples showed that the major tooth types of modern-day 525 Oreochromis are also represented in the fossil record. Their total morphological 526 variation is largely restricted to the morphospace occupied by contemporary 527 teeth, suggesting a close functional relationship between ancient populations 528 and their likely descendant, O. hunteri. 529

530 2.2 Introduction

The study of relatively young evolutionary radiations in geographically confined 531 areas has been the key to understanding mechanisms of speciation and adaptive 532 divergence. In this respect, cichlid fishes have provided one of the best 'natural 533 experiments' to study radiations and their prevalence in nature; hence, they 534 function as a model system to examine animal diversification and its underlying 535 processes (reviewed in Salzburger, 2009). The natural history of African cichlid 536 radiations is profoundly influenced by the tectonic and climate history of 537 the continent, and the East African Rift in particular (Koblmüller et al., 538 2008; Salzburger et al., 2014). Lake-level fluctuations during climatic dry-539 wet cycles and the development of the rift have alternatingly created and 540 eliminated freshwater habitats, and they changed the hydrographic connectivity 541

of rivers and lakes, thereby affecting gene flow between populations. Extrinsic
environmental factors, such as these tectonic and climatic events, are believed
to interact in complex ways with intrinsic biological characteristics to drive
speciation and evolutionary radiation (Koblmüller *et al.*, 2008; Salzburger *et al.*,
2014).

One trait responsible for the evolutionary success of cichlids is their 547 complex and extraordinarily adaptable trophic apparatus (e.g., Fryer & Iles, 548 1972). The oral and pharyngeal jaws and their associated dentition can change 549 over the course of a few generations (van Rijssel et al., 2015) or even within 550 the lifetime of a single specimen (Meyer, 1987; Gunter et al., 2013). Because 551 of this flexibility, cichlid fishes can exploit an extensive array of food resources, 552 and do so with impressively fine-scaled resource partitioning, especially in 553 rock-dwelling species (e.g., Fryer, 1957; Reinthal, 1990). As a result, disruptive 554 natural selection may lead to rapid speciation and the maintenance of distinct 555 species, even if differences in feeding ecology and occupied niches appear to be 556 modest (Schluter, 2000). 557

Current knowledge about the processes of diversification and speciation in 558 cichlids is often derived from evolutionary patterns revealed in the phylogenetic 559 trees of extant taxa (e.g., Danley & Kocher, 2001; Sato et al., 2003; Barluenga 560 et al., 2006; Wagner et al., 2012). With little concrete information on extinct 561 species, however, a large knowledge gap remains. In this regard, fossil data 562 extracted from African lake sediment records possess great potential. They may 563 increase our understanding of both biological evolution and long-term ecological 564 processes such as habitat specialisation and niche partitioning, provided that 565 these data are continuous across environmental transitions (Park & Cohen, 566 2011) and allow analysis at sufficiently high temporal resolution (Van Bocxlaer 567 & Hunt, 2013). Cichlid fish fossils preserved in sediments of Africa's large 568 tectonic lakes remain relatively unexplored in this respect (Reinthal et al., 569 2011), partly because a combination of great species richness, enormous habitat 570 diversity and complex biotic interactions (Fryer & Iles, 1972) hampers the 571 constraint of fossil data and the correct interpretation of fossil assemblages. 572

The sediment record of crater Lake Chala in East Africa (Kenya/Tanzania) contains abundant fossil teeth, bones and scales of cichlid fishes (Fig. 1.4). These remains accumulated throughout at least the last 25,000 years and have, hence, produced a unique system to study long-term ecological and evolutionary

dynamics. The lake's extant fish community consists of four cichlid species, of 577 which only one is indigenous and endemic: the Lake Chala Tilapia, Oreochromis 578 hunteri Günther, 1889 (Trewavas, 1983; Seegers et al., 2003). This is the fish 579 known locally as 'Chala', and here abbreviated as 'OH'. Three other species 580 have been introduced during the second half of the 20^{th} century: the Korogwe 581 tilapia, O. cf.korogwe Lowe, 1955 ('Bandia', here abbreviated as 'OK'); the 582 redbreast tilapia, Coptodon rendalli Boulenger, 1896 ('Kijiji') and a small 583 haplochromine cichlid referred to as *Haplochromis* spec. "Chala" by Seegers 584 et al. (2003). Therefore, the only cichlid remains we expect to find in the 585 pre- 20^{th} century sediment record are those of the endemic *O. hunteri*. Other 586 fish remains may belong to other indigenous fishes that have become (locally) 587 extinct since. 588

Lake Chala is the focus of multidisciplinary investigations into the long-589 term climate and environmental history of equatorial East-Africa, and these 590 provide a highly suitable context for detailed study of the lake's fossil record. 591 A wide array of paleo-environmental proxies extracted from its sediments 592 have allowed reconstruction of lake-level fluctuations (Moernaut et al., 2010), 593 temperature change (Sinninghe Damsté et al., 2012) and diverse other aspects 594 of the lake's environmental history, which facilitate linking long-term aquatic 595 ecosystem dynamics with pacers of environmental change. 596

Although the continuous sediment record of Lake Chala creates interesting 597 opportunities to study evolutionary and ecological changes through time, pre-598 served fossil fish remains mainly consist of individual teeth. Notwithstanding 599 the vast amount of descriptive literature on trophic morphology and its eco-600 logical significance (e.g. Greenwood & Gee, 1969; Witte & Van Oijen, 1990), 601 contemporary ecomorphological studies focus mainly on morphometric analyses 602 of complete oral and pharyngeal jaws (see e.g. Kerschbaumer & Sturmbauer, 603 2011). A knowledge gap thus remains, concerning the application of morpho-604 metric methods on individual teeth with the objective to infer feeding ecology 605 in cichlids. 606

In this Chapter, we develop a methodology to quantify variation in the morphology of oral teeth using geometric morphometrics. These efforts are a first step towards the ecological or evolutionary interpretation of changes in the morphology of fossil teeth through time. To build a reference framework for the interpretation of fossil data, we assessed different components of the overall

variability in tooth shape present in the two Oreochromis species inhabiting Lake 612 Chala today. First we used landmark analyses to document variation in overall 613 body morphology of the endemic O. hunteri and introduced O. cf.korogwe, 614 with the aim to unambiguously differentiate individuals of both taxa. Using 615 representative specimens of each species, we then performed traditional length 616 measurements on complete oral teeth and geometric-morphometric analyses of 617 tooth crowns. We also compared the performance of semi-landmark analysis 618 (SLM) and Elliptic Fourier Analysis (EFA) in the quantitative characterisation 619 of tooth crown morphology. Our results allowed us to draw tentative inferences 620 about the feeding ecology of both species. Finally, we compared the combined 621 morphospace occupied by oral teeth of the extant Lake Chala fishes with those 622 found in three fossil assemblages of up to 24,000 years old, as a first exploration 623 of morphological variation through time. 624

⁶²⁵ 2.3 Materials and methods

626 2.3.1 Specimen collection and preparation

627 Modern material

A total of 75 specimens of O. hunteri (OH) and 20 of O. cf. korogwe (OK) were 628 used to examine overall body morphology (Fig. 2.1). Freshly caught specimens 629 were identified by local fishermen and positioned on graph paper with their 630 left side up and fins spread out. They were photographed from a standard 631 angle (i.e. perpendicular to the dorsoventral and anteroposterior axis) using 632 a Canon EOS400D SLR camera with an 18-55 mm lens. The photographs 633 were pre-processed in Photoshop CS5 to increase contrast, and subsequently 634 stacked in a .tps file using tpsUtil version 1.58 (Rohlf, 2013c) for morphometric 635 analyses. 636

Gualitative microscopic examination of the trophic apparatus of both Oreochromis species revealed great variation in the morphology of oral teeth within a single specimen, compared to their overall variation between specimens of the same species or even between the two species. Therefore, we performed an exhaustive study of all teeth in the left half of the oral jaws, of one representative specimen of each species. Visual inspection of the pharyngeal apparatus showed a large number (ca. 1000) of slender ('papilliform') unicuspid teeth in both

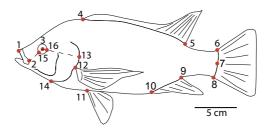


Figure 2.1: Outline drawing of a female *Oreochromis hunteri* from Lake Chala with indication of the digitised landmarks. 1 rostral tip of the upper jaw, 2 caudo-dorsal tip of maxillary bone, 3 centre of the eye, 4 rostral insertion of the dorsal fin, 5 caudal insertion of the dorsal fin, 6 base of the dorsal caudal fin ray, 7 intersection between lateral line and insertion of the caudal fin, 8 base of the ventral caudal fin ray, 9 caudal insertion of the anal fin, 10 rostral insertion of the anal fin, 11 rostral insertion of the pelvic fin, 12 base of the dorsal pectoral fin ray, 13 most caudal point of the operculum, 14 ventral intersection between the branchiostegal membrane and body outline, 15 intersection between the line connecting landmarks 2 and 3 and the eye outline

species. Some preliminary analyses were undertaken on these pharyngeal teeth,but are not discussed here.

After fixation of fish heads in formalin 35%, dentaries and premaxillae 646 were dissected and stained using an Alizarin red S solution, and soft tissues 647 were removed by maceration in 0.5% KOH following Wautier *et al.* (2002). 648 Because each oral jawbone contains ca. 100-160 teeth and because of bilateral 649 symmetry, all data presented are from teeth in the left dentary (DTL) and 650 premaxilla (PML). Overview photographs of each jawbone were taken to keep 651 track of the exact location of individual teeth (Fig. 2.2), which were divided 652 into five groups. The first group (DTL1/PML1) comprises the outer series of 653 oral teeth. Group DTL2/PML2 constitutes the outer row of the inner series 654 of teeth, from here on referred to as the second series, which is similarly well 655 defined. The other teeth are more chaotically positioned and not evidently 656 arranged in series. Consequently, we split them in an anterior (DTLA/PMLA), 657 middle (DTLM/PMLM) and posterior (DTLP/PMLP) group. Soft tissue 658 maceration for 20-25 h allowed easy extraction of all teeth from the jawbones, 659 under a Nikon SMZ-U binocular microscope (zoom 1:10). 660

Each tooth was mounted in glycerin on a microscope slide and oriented with its convex side up. Photographs were taken at 60x magnification, using an Olympus Camedia C5060 Wide Zoom, which is a 5.1 megapixel camera with HQ lens and 4x optical zoom, attached to the optic mount of the binocular microscope. Because the strong curvature of large teeth hampered standardised orientation, their base was clipped with a scalpel in such a way that the crown could be mounted in a flat plane. After imaging, all teeth were stored individually in 100% ethanol.

The contrast of tooth images was enhanced in Photoshop CS5. Pho-669 tographs of PML teeth were mirrored to allow direct comparison with DTL 670 teeth, because teeth in the premaxilla develop as counterparts to those in the 671 dentary. For consistency, also tricuspid teeth were mirrored, although they are 672 bilaterally symmetrical and mirroring does not generally affect their morphol-673 ogy. Photographs of individual teeth were combined in a .tps file using tpsUtil 674 and served as input files for length measurements and geometric-morphometric 675 digitisation. 676

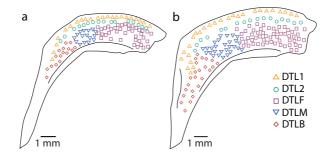


Figure 2.2: Left dentaries (dorsal view) of *Oreochromis hunteri* (a) and *O. cf. korogwe* (b) showing the position of individual oral teeth, subdivided into five groups: outer series (DTL1), second (outer row of inner) series (DTL2), and the anterior (DTLA), middle (DTLM) and posterior (DTLP) portions of the other inner series

677 Fossil material

In 2003 and 2005, the CHALLACEA project recovered a 21.65-m-long sequence of mostly finely laminated organic muds from the middle of Lake Chala (Verschuren *et al.*, 2009). Excision of five turbidite sections yielded a 20.60-m-long sequence of continuous offshore lacustrine sedimentation spanning the last 25,000 years (Blaauw *et al.*, 2011). Four large-volume sediment samples (Table

683 2.1), each representing a ca.100-year time interval, had been extracted partly

from the master core sequence and partly from overlapping core sections, origi-684 nally for the purpose of compound-specific radiocarbon dating. For this pilot 685 study, these bulk samples were subsequently processed by adding 10% NaOH 686 and heating to 75°C for ca. 30 min, while stirring continuously. Subsequently, 687 the samples were sieved on stacked 250- and $150-\mu m$ mesh sieves to recover 688 all particles larger than 150 μ m in two size fractions. Both fractions were 689 searched using a Nikon SMZ-U Zoom 1:10 binocular microscope. All fossil 690 teeth, oral and pharyngeal, were picked with a brush, photographed and stored 691 in Eppendorf vials. The sample dated to 19,100 year BP did not contain any 692 fossil fish teeth. Measurement and morphological analysis of fossil teeth was 693 identical to that of the modern teeth. Fossil bicuspid teeth derived from right 694 dentaries and left premaxillae were mirrored, for direct comparison with the 695 left dentary and right premaxilla. Given the bilaterally symmetric nature of 696 tricuspid teeth, it was impossible to determine which lateral side they derived 697 from, and hence no mirroring of fossil tricuspid teeth was performed. Teeth 698 from pharyngeal and oral jaws have strongly different morphologies, so that 699 they can be assigned unambiguously to the correct jaw. 700

Table 2.1: Sediment samples from Lake Chala screened for recovery of fossil fish teeth. Samples are indicated with their depth in the sequence, calibrated age range, sample size (estimated dry sediment weight) and the number of oral teeth recovered; any pharyngeal teeth found are not included. The oldest sample originates from a turbidite section later excised from the composite sediment sequence, and hence its approximate age (ca. 500 years older than adjacent muds; Blaauw *et al.*, 2011) and identical top and basal depths

Depth top (cm)	Depth base (cm)	Age range (cal yr BP)	Estimated dry weight (g)	# fossil oral teeth
572	581	7,000-7,100	17.83	25
666	674	8,200-8,300	21.02	15
1520	1527	19,100-19,200	91.22	0
1951	1951	ca. 24,000	60.62	11

701 2.3.2 Measurements and morphometric analyses

702 Overall body morphology

Overall body morphology was documented largely following Kerschbaumer &
Sturmbauer (2011) with sixteen traditional landmarks (Fig. 2.1) in tpsDig2
v.2.17 (Rohlf, 2013a). Size was calibrated using the graph paper on which

specimens were placed (see above). The digitised landmark dataset was aligned 706 via Procrustes superimposition in the program CoordGen6h of the Integrated 707 Morphometrics Package software (IMP; Sheets, 2008). After all variation in 708 size, position and orientation was removed. Procrustes coordinates and rescaled 709 centroid size were saved as a data matrix file in IMP format. Centroid size is 710 a reliable size measure (e.g., Bookstein, 1997), and other fish studies report 711 strong correlations between centroid size and traditional length measurements 712 (Langerhans & DeWitt, 2004; Leinonen et al., 2006). The overall variation in 713 body shape was then analysed via principal component analysis (PCA) on the 714 Procrustes superimposed coordinates. We performed model-based clustering on 715 the first two principal components (PC) without a priori species designation, 716 to test whether OH and OK can be reliably differentiated based on body shape. 717 These analyses were performed in 'Mclust' v.4.3 (Fraley et al., 2012) in R (R 718 Development Core Team, 2016), which was used for all statistical analyses. The 719 performance of different clustering algorithms was evaluated using a Bayesian 720 Information Criterion (BIC), which documents the optimal number of clusters 721 that can be assigned to the data without over-fitting. We also determined 722 the axes of maximal group separation using discriminant function analysis 723 (DFA). A partial warp matrix was generated from the Procrustes superimposed 724 coordinates using tpsRelw v.1.53 (Rohlf, 2013b), and we performed a DFA 725 on these partial warp scores in PAST v.2.17 (Hammer et al., 2001), in order 726 to reveal more subtle morphological differences between the species. Finally, 727 because the analysed OH specimens were generally larger than OK specimens, 728 we studied the extent to which variation is caused by growth allometry. First, 729 we performed a general least-squares regression of PC1 versus log centroid 730 size (LCS), and subsequently we executed a MANCOVA on PC1 and LCS 731 using tpsRegr version 1.40 (Rohlf, 2011) to compare slopes and intercepts 732 between both species. If both nominal species would display similar slopes 733 and intercepts, they would share the same relationship between shape and size 734 (LCS), which would suggest that all variation is caused by allometry. This 735 scenario would then imply that we had collected a biased sample of one species. 736 rather than of the two morphologically distinct species identified by Trewavas 737 (1983) and Seegers *et al.* (2003). 738

739 Tooth measurements

Traditional length measurements on complete teeth were used to evaluate 740 possible differences in tooth shape between the two nominal Oreochromis 741 species, parallelling overall differences in body morphology. Although these 742 measurements are less time-consuming than geometric-morphometric analysis 743 of tooth crown morphology, they cannot readily be used for measuring fossil 744 teeth, because these are often broken in ways that do not allow reconstructing 745 general tooth morphology. Moreover, the measurements may reveal information 746 that is complementary to that obtained via geometric morphometrics. Length 747 and width of complete, unclipped teeth were measured in number of pixels, 748 using the measurement tool in tpsDig2. Length was measured along the median 749 axis of the tooth neck as the distance from the top of the major cusp to the 750 tooth base (Fig. 2.3); width was measured at the narrowest point of the neck, 751 and subsequently the length/width ratio was calculated. We performed a 752 non-parametric KruskalWallis Rank test with these ratios to discern differences 753 between the two species. 754

755 Tooth digitisation for geometric morphometrics

We examined crown morphology in more detail with geometric-morphometric 756 methods. To ascertain the most suitable method for our data, we compared 757 two common methods, semi-landmark (SLM) and elliptic Fourier analysis 758 (EFA). For SLM, a number of homologous points must be established, which 759 can be supplemented with open curves. We positioned landmarks one and 760 three (LM1 and LM3) at the base of the crown, on both sides where the 761 enameloid-dentin junction reaches the tooth outline (Fig. 2.3). Landmark 762 two (LM2) was positioned at the top of the major/middle cusp, by shifting 763 the baseline defined by LM1 and 3 until it lies tangential to the most apical 764 point of the major cusp. If, due to wear or abrasion, the location of these 765 landmarks was not clear, the tooth was excluded from analysis. Eventually, we 766 analysed 468 teeth, 51 of which were fossils. Although the homology criterion 767 is somewhat relaxed for LM2, inclusion of this landmark significantly increased 768 alignment accuracy compared to analyses that only used LM1 and LM3 (C. 769 Manntschke, unpublished data). Two open curves, each with 20 equidistant 770 semi-landmark points digitised in clockwise direction, were anchored between 771

LM1 and LM2, and between LM2 and LM3. Digitisations for SLM were 772 performed in tpsDig2. The generated semi-landmarks were initially specified 773 as landmarks using tpsUtil, so that partial Procrustes super-impositioning 774 could be performed in CoordGen6h. Generated Procrustes coordinates and 775 centroid size were saved as a data matrix file in IMP format. Centroid sizes 776 are comparable between teeth without additional calibration, because all teeth 777 were imaged at the same magnification (60x). These data were subsequently 778 loaded in SemiLand6 to align the semi-landmarks by sliding them along the 779 curves using the perpendicular projection criterion (Perez et al., 2006). 780

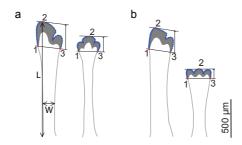


Figure 2.3: Outline drawings of representative bicuspid and tricuspid oral teeth from *Oreochromis hunteri* (a) and *O.* cf. *korogwe* (b), with enameloid-covered area shaded grey. Semi-landmark analysis (SLM) involved three landmarks (1-2-3), supplemented with two open curves consisting of twenty equally spaced semilandmarks

EFA was performed using a closed curve tracing the outline of the tooth 781 crown and connecting the LM1 and LM3 positions on the enameloid-dentin 782 boundary. Before outlines were generated, tooth images were converted into 783 black and white bitmap files with Photoshop CS5 to maximise contrast. Further 784 analyses were carried out using the SHAPE software package (Iwata & Ukai, 785 2002). All generated bitmap files were loaded into SHAPE ChainCoder v.1.3, 786 where outlines were traced in counter-clockwise direction, and a chain code file 787 was generated. Contour data were aligned automatically, and the digitised chain 788 codes were imported in SHAPE CHC2NEF v.1.3a to obtain Fourier coefficients. 789 The maximum number of harmonics was set to 50, and standardisation for 790 size, rotation and starting point was based on the first harmonic. 791

792 Comparison of SLM and EFA

For comparison of SLM and EFA, data were ordinated using nonmetric multi-793 dimensional scaling (NMDS) based on Procrustes superimposed coordinates 794 (SLM) and Fourier coefficients (EFA). We used NMDS because of its limited 795 assumptions, because the number of components can be chosen and because the 796 analysis generates a stress value (i.e. goodness of fit), which makes comparison 797 between the methods more straightforward. We constructed Euclidean dis-798 tance matrices for both datasets and performed ordination in two dimensions. 799 Preliminary runs reached convergence before 20 iterations, and hence we set 800 the maximum number of iterations to 50 and performed the analyses with 801 1000 random starting configurations. Stress values were obtained, multiplied 802 by 100 and evaluated following Kruskal & Clarke's rules of thumb (Kruskal, 803 1964; Clarke, 1993), indicating ordination to be good if stress values are <10. 804 Analyses were performed in R, using the packages 'vegan' (v.2.0.10; Oksanen 805 et al., 2013) and 'MASS' (v.7.3.34; Venables & Ripley, 2002). Furthermore, we 806 analysed the correlation of pairwise distances in order to quantify the similarity 807 in how both methods handle tooth shape variation. We performed Mantel 808 tests on Euclidean distance matrices constructed on the basis of superimposed 809 coordinates (SLM) and Fourier coefficients (EFA); these permutate rows and 810 columns randomly and yield distribution matrix correlations for unrelated 811 matrices (Dietz, 1983). We used the package 'ade4' (v.1.6.2; Chessel et al., 812 2004) with 1,000 random permutations. 813

⁸¹⁴ Comparison of tooth morphology in modern and fossil fishes

Morphospace occupation of fossil and modern fish teeth was assessed with 815 PCA. To check whether our *a priori* classification of bi- and tricuspid teeth 816 was consistent with their location in morphospace, we performed model-based 817 clustering with 2 models on the first 2 PCs and examined the error rate. 818 Subsequently, the first principal component score was analysed and compared 819 between groups to assess the significance of any differences in average tooth 820 morphology between the fossil and the modern-day Oreochromis populations 821 in Lake Chala. We therefore performed a Kruskal-Wallis test in R, for bicuspid 822 and tricuspid teeth separately, followed by *post-hoc* comparisons using Tukey 823 HSD tests (included in the standard package of R). 824

825 2.4 Results

826 2.4.1 Overall body morphology

Ordination (PCA) of the landmark coordinates of 75 O. hunteri (OH) and 20 827 O. cf. koroqwe (OK) specimens from Lake Chala shows a strongly bimodal 828 pattern in overall body shape along the first principal component (PC1; Fig. 829 2.4a). Main sources of variation on PC1 reflect the length of the tail (long for 830 OH, short for OK) and body depth (slender, shallow habitus for OH; more 831 robust, deeper body for OK). Variation on PC2 mainly reflects differences in 832 the shape of the head, as represented by the relative distance between the 833 rostral tip of the upper jaw and the most caudal point of the operculum (Fig. 834 2.1). Applying model-based clustering to the first two principal components 835 yielded greatest support for solutions with only two clusters ($\Delta BIC \geq 5.5$ 836 over models with fewer or more clusters). These two clusters coincide entirely 837 with the two species as identified in the field, apart from one aberrant, slender 838 OK specimen that clusters with OH. These results are further corroborated 839 by the DFA, which separates groups along a highly similar morphological 840 axis. Regression of PC1 versus centroid size indicates a strong and significant 841 relationship between specimen size and shape, and hence allometric variation 842 (t-value = 11.45, p < 0.001). MANCOVA results indicated that OH and OK 843 do not differ significantly in the slope of allometric growth (Fig. 2.4b: Wilks' 844 K = 0.64, p = 0.99), but they do in intercepts (Wilks' K = 0.071, p < 0.001), 845 which suggests that the relationship between size and shape differs between the 846 two species, supporting the taxonomic division between OH and OK. Hence, 847 the shape differences between both species do not solely result from allometry. 848

849 2.4.2 Shape of oral teeth

Qualitative visual inspection of tooth shape and arrangement in the oral jaw 850 bones (dentary and premaxilla) hints towards a strong similarity between the 851 two studied Oreochromis species (Fig. A.1). In both species, the outer series 852 (DTL1/PML1; Fig. 2.2) consists of bicuspid teeth, whereas the second series 853 (DTL2/ PML2) and more inward placed teeth are mostly tricuspid, except 854 for some of the posterior teeth (DTLP/ PMLP) which are more unicuspid in 855 shape (but distinct from unicuspid pharyngeal teeth), mostly in the dentaries. 856 No obvious differences were observed between teeth from the upper (PML) and 857

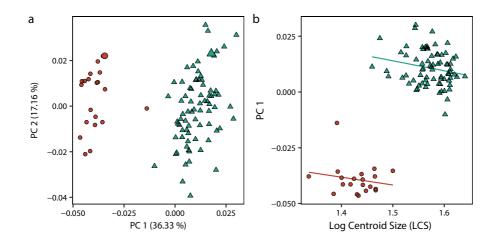


Figure 2.4: a PCA ordination plot showing the variation of overall body morphology in modern-day Lake Chala Oreochromis, based on 16 landmarks (cf. Fig. 2.1) measured on 75 O. hunteri (green △) and 20 O. cf.korogwe (red ○) specimens. The two enlarged symbols represent the specimens used for analysis of tooth morphology. b MANCOVA test on overall body morphology of both species corrected for size, showing similar regression slopes of growth allometry but significantly different intercepts

lower (DTL) jaws for either species. The total number of oral teeth differed considerably between OH (n = 268) and OK (n = 307), but examination of more specimens is required to examine whether this difference is consistent. However, the number of oral teeth that could be included in the geometric morphometric dataset is similar (Supplementary Table A.1).

Traditional length measurements revealed that the length/width ratio of 863 oral teeth of both species is significantly different (Kruskal-Wallis $\chi^2 = 116.3$, 864 p < 0.001; Fig. 2.5). The range in tooth length is roughly equal for both 865 species, but OK teeth are wider at their narrowest point. The difference is most 866 pronounced in the longest teeth. Shorter teeth show a clear overlap between 867 the species, but nevertheless the teeth of OK can be considered more stout 868 than those of OH. Teeth with overlapping morphology mostly derive from the 869 inner series (DTL/ PML A-M-P). 870

Model-based clustering on the first two principal components classified six out of 468 teeth differently (1% error rate), and hence both methods of clustering are considered useful. For further analyses, we classified specimens following the generated model.

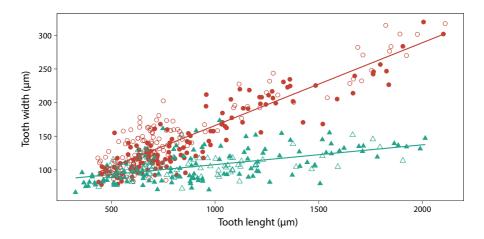


Figure 2.5: Length/width ratio of oral teeth from modern-day *Oreochromis hunteri* (green \triangle ; n = 212) and *O*. cf.*korogwe* (red \circ ; n = 259), distinguishing between teeth from dentaries (hollow symbols) and premaxillae (filled symbols). Best-fit linear regression lines are shown per species. Length is the distance from the top of the major cusp to the base of implantation, and width is measured at the narrowest point of the neck (cf. Fig 2.3)

⁸⁷⁵ 2.4.3 Comparative analysis of SLM and EFA

The stress values associated with NMDS were fairly low for both SLM (8.17) and 876 EFA (10.13), indicating that the data can be fit well in two dimensions, and that 877 NMDS ordination plots (Fig. 2.6) can be considered a reliable representation 878 of the variation in the data. Both methods reveal a marked separation between 879 bicuspid and tricuspid teeth along NMDS1. The few unicuspids present in the 880 DTLP/PMLP group form an elongated 'tail' of NMDS scores on the negative 881 end of NMDS2 in both graphs. Most importantly, the position of the five groups 882 is similar for both methods, suggesting that they capture the overall variability 883 in tooth crown morphology in a very comparable manner. This conclusion is 884 supported by a Mantel test between the Euclidean distance matrices of the 885 two datasets, which reveals a strong and highly significant positive correlation 886 (z = 0.81, p < 0.001).887

2.4.4 Tooth crown morphology of extant Oreochromis

Ordination based on the SLM coordinates of all modern tooth crowns resulted in a major separation between outer-series (mainly bicuspid) teeth (DTL1/

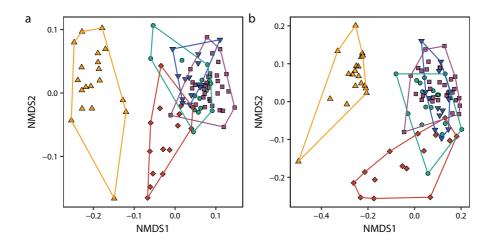


Figure 2.6: Nonmetric multidimensional scaling (NMDS) plots for the crown morphology of mandibular oral teeth in *O. cf.korogwe*, assessed using semi-landmark analysis (SLM; a) and elliptic Fourier analysis (EFA; b). Symbols represent the five pre-defined groups displayed in Fig. 2.2: First (yellow △) and second series teeth (green ○), and anterior (purple □), middle (blue ∇) and posterior (red ◊) inner teeth.

PML1: PC1-values < -0.1) and inner-series (mainly tricuspid) teeth (other 891 groups: PC1-values > -0.1; Fig. 2.7a). Crowns with a somewhat aberrant 892 unicuspid shape plotted mostly on the negative side of PC2 and display some 893 overlap with tricuspid teeth. No obvious differences were observed between 894 DTL and PML teeth belonging to the same group, suggesting that, within an 895 individual specimen, the upper and lower jaw bones bear highly similar teeth. 896 Notably however, both species have several unicuspid teeth in the posterior 897 part of the dentary, but not in the premaxilla (Supplementary Figure A.3). 898 For each tooth type, teeth of the two species occupy largely distinct regions in 899 morphospace, separated along PC2 (Fig. 2.7a). 900

Because large differences between the major tooth types may obscure more 901 subtle differences of tooth crown variation, outer-series and inner-series teeth 902 were subjected to separate ordinations. After omitting two highly aberrant 903 specimens to enhance spread of datapoints, the PCA of outer-series teeth (Fig. 904 2.7b) shows that the separation between OH and OK is mainly driven by 905 differences in the relative size of the major crown cusp: bicuspid teeth of OH 906 tend to have a large major cusp, and hence more positive PC1 values, whereas 907 bicuspid teeth of OK, with a relatively smaller major cusp, have negative 908

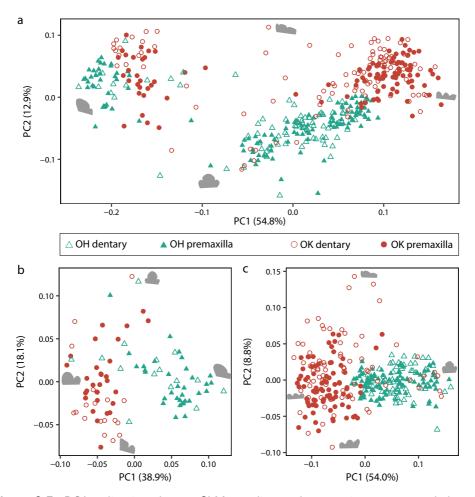


Figure 2.7: PCA ordination plots on SLM coordinates documenting crown morphology in Oreochromis hunteri (green △) and O. cf. korogwe (red ∘), distinguishing between teeth from dentaries (hollow symbols) and premaxillae (filled symbols). a All modernday teeth analysed; b outer-series (mostly bicuspid) teeth only; c inner-series (mostly tricuspid) teeth only. Silhouettes of some representative crown shapes facilitate interpretation of the morphological variation along PC1 and PC2

PC1 values. Variation on PC2 seems to primarily reflect differences in the occlusion surface. OK bicuspids tend to have a sharp major cusp (positive PC2 values), whereas those of OH have a convex, more broadly rounded major cusp (negative PC2 values). Aberrant teeth are also singled out in morphospace, for example, the small number of unicuspid DTLP/PMLP teeth (located between the clusters, but slightly closer to the bicuspids cluster). In addition, one outer-series tooth of the analysed OH specimen had a tricuspid, rather than ⁹¹⁶ the usual bicuspid, appearance.

2

Ordination of inner-series teeth separately (Fig. 2.7c) shows that both 917 taxa are well-differentiated on PC1, which mainly captures variation in the 918 relative size of the central cusp: negative values indicate a small central cusp, 919 as typically found in OK, whereas positive values indicate a large central cusp. 920 as found in OH. PC2 captures information related to asymmetry, a biological 921 signal that is largely related to tooth wear and, hence, rather to function 922 than to original morphology. Original morphology and tooth wear are easily 923 distinguished from one another based on the enameloid-dentin boundary: a 924 sloping boundary suggests an originally asymmetric morphology, whereas a 925 level boundary points towards asymmetry after wear. Highly negative PC2-926 values indicate that the right side of the crown is worn down, whereas crowns 927 with the left side worn down have highly positive PC2-values. Most crowns. 928 however, have PC2-values close to zero, and are more or less symmetrical. 929 attesting to rather limited wear. The range of PC1-scores is larger for OH 930 than for OK, indicating greater variation in the relative size of the middle cusp 931 in OH, whereas OK showed greater variation along PC2, indicating a greater 932 degree of asymmetry caused by wear. Teeth from the inner series (DTL/PML 933 A-M-P) of the two species overlap partly in crown shape. 934

⁹³⁵ 2.4.5 Teeth from the fossil archive of Lake Chala

The crowns of 51 fossil teeth from Lake Chala were analysed, and these teeth 936 occupy a combined morphospace that is very similar to that of modern-day 937 teeth (Fig. 2.8), demonstrating that the morphology of fossil teeth is similar to 938 those of the extant *Oreochromis* inhabiting Lake Chala. Hence, cichlid fishes 939 with bicuspid and tricuspid teeth have been present in Lake Chala over at 940 least the past 24,000 years. Nevertheless, some fossil teeth plot outside the 941 morphospace bounds of the major modern-day teeth types, indicating that 942 their morphology is slightly unusual. 943

The elongated cluster with largely positive PC1-values represents tricuspid teeth. The 10 fossil tricuspid teeth dated to 7,000 years ago (year BP) are spread throughout this cluster, but do not exceed the morphospace boundary of modern inner-series teeth. Average crown morphology was thus similar to that observed within the two examined extant specimens, and no notably different oral tooth shapes were encountered in the fossil record. Roughly, three clusters

Table 2.2: Mean PC1 of bicuspid and tricuspid teeth in extant <i>Oreochromis hunte</i>	eri
(OH), extant O. cf. korogwe (OK) and the three samples of fossil teeth, and adjust	ed
p-values of post-hoc comparisons using the Tukey HSD test. P-values indicated with	an
* highlight significant results. The table illustrates that fossil tricuspid teeth resemb	ole
those of the native OH more closely in morphology than those of the introduced OI	K

	Mean PC1 \pm SD	Adjusted P values of Tukey HSD test				
		OH	OK	7,000 yr BP	8,200 yr BP	
Bicuspid teeth						
OH	-0.19 ± 0.04					
OK	-0.15 ± 0.04	$< 0.001^{*}$				
7,000 yr BP	-0.18 ± 0.02	0.96	0.04^{*}			
$8,200 \mathrm{~yr~BP}$	-0.15 ± 0.04	0.52	1	0.72		
24,000 yr BP	-0.18 ± 0.08	0.99	0.52	1	0.83	
Tricuspid teeth	ļ					
OH	0.03 ± 0.04					
OK	0.10 ± 0.04	$< 0.001^{*}$				
7,000 yr BP	0.05 ± 0.06	0.57	0.02^{*}			
$8,200 \mathrm{~yr~BP}$	0.00 ± 0.08	0.32	$< 0.001^{*}$	0.10		
24,000 yr BP	-0.02 ± 0.10	0.10	$< 0.001^{*}$	0.03^{*}	0.90	

can be distinguished among tricuspids from this time window. One cluster has 950 a crown shape similar to OK-tricuspids, one cluster resembles the OH tricuspids 951 most similar to those of OK, and a third cluster, with negative PC1-values, has 952 typical OH-crowns with a pronounced middle cusp. Almost all tricuspids from 953 this time window are located close to the longitudinal axis of the point cluster. 954 indicating great similarity and limited wear (compare with Fig. 2.7c). The 13 955 fossil tricuspid teeth dated to 8,200 year BP are distributed in a similar way. The 956 same three clusters (of roughly equal size) can be distinguished, but overall 957 they are displaced slightly towards more negative PC1-values compared to 958 those dated to 7,000 year BP, indicating that they possess a highly pronounced 959 middle cusp, similar to that of some outlier tricuspids in extant OH (cf. Fig. 960 2.7C). The 7 tricuspid teeth dated to 24,000 year BP show only two distinct 961 crown morphologies: one that is intermediate between the tricuspids of OH and 962 OK, and one with an extremely pronounced middle cusp, sometimes even more 963 pronounced than in outlier modern-day OH tricuspids. In general, the average 964 crown morphology of the tricuspid teeth migrates over time (from 24,000, 8,200 965 and 7,000 year BP to the modern fauna) towards more positive values on PC1. 966

This shift is partly corroborated by the significance of a Kruskal-Wallis test 967 on the differences between mean PC1-scores (Kruskal-Wallis $\chi^2 = 156.3$, p 968 $< 2.2 \times 10^{-16}$). Post-hoc comparisons using the Tukey HSD test (Table 2.2) 969 indicate that modern-day OH and OK tricuspids differ significantly in the 970 average pronunciation of their middle cusp. All fossil tricuspid groups differ 971 significantly from OK, and hence more closely resemble the native OH than 972 the introduced OK. Between fossil groups, tricuspids from 24,000 year BP have 973 a significantly different morphology than the ones from 7,000 year BP. 974

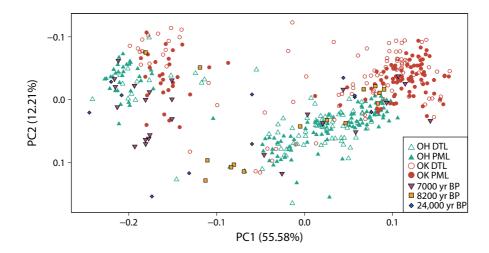


Figure 2.8: PCA ordination of SLM-coordinates documenting crown morphology of all modern-day teeth, and fossil teeth of the three age intervals (7,000, 8,200 and 24,000 years ago) analysed

Fossil bicuspid teeth all plot within or near the morphospace occupied by 975 modern bicuspids, with negative PC1-values. The 15 fossil bicuspid teeth dated 976 to 7,000 year BP span the variation seen in modern-day OH and OK, and some 977 specimens with negative PC2-scores plot outside the modern morphospace 978 occupation (Fig. 8). Bicuspids from sediments dated to 8,200 year BP (n = 2)979 fall within the crown variability found in OK, whereas those dated to 24,000 980 year BP (n = 4) more closely resemble bicuspids of modern-day OH. Post-hoc 981 comparisons, however, indicate that only bicuspids from 7,000 year BP are 982 significantly different from modern-day OK bicuspids; none of the three groups 983 of fossil bicuspid teeth differ significantly from one another (Table 2.2). 984

985 2.5 Discussion

⁹⁸⁶ 2.5.1 Comparative analysis of SLM and EFA

The EFA and SLM methods both have merits that make them useful tools 987 to document tooth crown morphology. The main advantage of EFA is that 988 this method is more robust against violations of the homology assumption 989 (Van Bocxlaer & Schultheiß, 2010). At the downside, closing the crown outline 990 with a straight line along the (poorly-defined) crown-neck boundary results in 991 propagation of error along this line, because points along it are not independent. 992 This creates artificial shape variation which potentially blurs other, more 993 biologically relevant shape information. However, difficulty to accurately define 994 the crown-neck boundary implies that locating LM1 and LM3 for SLM was not 995 always unambiguous either. Furthermore, our method to locate LM2 directly 996 depends on the positions of LM1 and LM3, so that some uncertainty may have 997 been propagated towards LM2 as well. Because both methods produced very 998 similar and fully consistent results, and because landmark-based approaches are 999 conceptually more straightforward, we adopted SLM as the principal method 1000 in this study. Moreover, SLM data are more directly comparable to those of 1001 the landmark analysis used to characterise overall body morphology. 1002

2.5.2 The modern-day *Oreochromis* of Lake Chala: morphol ogy and feeding ecology

The clearly bimodal distribution of overall body morphology, based on 16 1005 landmark characters, of modern-day *Oreochromis* in Lake Chala supports the 1006 presence of two species, O. hunteri and O. cf. koroque (Trewavas, 1983), which 1007 are also recognised by local fishermen as morphologically distinct. Body depth 1008 accounts for most of the shape variation between the two species. The lack of 1009 overlap and limited number of intermediate forms supports the taxonomic status 1010 of these species and argues against the occurrence of significant interbreeding 1011 between both species in Lake Chala. One slender OK specimen clusters 1012 with OH and potentially represents a rare hybrid specimen. However, the 1013 mitochondrial DNA of this specimen (BANDIA cha022) is identical to other 1014 O. cf. koroque specimens (Chapter 3), arguing against its hybrid status. The 1015 natural distribution of O. koroque is limited to the lower parts of the Pangani 1016 River basin in northeastern Tanzania (Trewavas, 1983). It was introduced in 1017

Lake Chala (Seegers *et al.*, 2003), most likely in the 1950s or early 1960s, by a colonial-era owner (Caxton Oluseno, pers. comm. 2014). Hence, we do not expect to find remains of this species in the lake's (pre- 20^{th} century) fossil record.

The ecology and feeding habits of O. hunteri and O. cf. korogwe are 1022 poorly known. A concise description in Trewavas (1983) suggests that the 1023 young of *O. hunteri* feed on algae and organic debris, but no information on 1024 adults is given. The diet of O. korogwe is unknown, except that in fish-culture 1025 ponds the species feeds on algae and insects. This information clearly does not 1026 suffice to draw conclusions on their respective diets in Lake Chala, whereas 1027 our biometric and geometric-morphometric data on the oral jaw teeth of both 1028 species do allow tentative inferences on feeding ecology. We revealed substantial 1029 morphological differences between OH and OK in overall body shape as well as 1030 in the morphology of their oral teeth, both general and crown-specific. Given 1031 these recurrent and consistent differences, it is likely that the two Oreochromis 1032 species inhabiting Lake Chala exploit different ecological niches. 1033

The large number of oral teeth in wide, U-shaped jaws, and the combination 1034 of an outer series of bicuspid teeth with multiple inner series of tricuspid teeth 1035 (as well as a large number of papilliform pharyngeal teeth, not studied in detail 1036 here) suggest that both O. hunteri and O. cf. koroque graze on epilithic algae 1037 (e.g., Witte & Van Oijen, 1990; Rüber & Adams, 2001). In the large African rift 1038 lakes, such epilithic grazers can form ecologically stable cichlid communities with 1039 very high species diversity (Reinthal, 1990). Notwithstanding that multiple 1040 species may all appear to feed on highly similar resources, namely algae 1041 growing on rocks, fine-scaled resource partitioning within such communities 1042 has been described extensively (e.g., Fryer, 1957; Genner et al., 1999), and 1043 this partitioning has been linked to an equally great diversity of trophic 1044 morphologies (Greenwood & Gee, 1969). The consistent and marked differences 1045 in tooth morphology of OH and OK, as revealed by this study, suggests that 1046 co-occurrence of the native and introduced *Oreochromis* in Lake Chala is 1047 facilitated by diet-related ecological differentiation. 1048

The widely separated, mobile and slender oral teeth of *O. hunteri* (Fig. 2.3a; Supplementary Fig. A.1) seem most appropriate for a diet of the socalled 'Aufwuchs', i.e. the organisms that live loosely attached within epilithic algae (Fryer, 1957). This Aufwuchs may include diatoms, other microscopic autotrophic organisms, microcrustacea and occasionally the nymphs and larvae
of aquatic insects. The slender shape of the teeth and their widespaced
arrangement allows them to function both as a comb and as a sieve. Moreover,
the bicuspid and second-series tricuspid teeth of OH have a relatively larger
and more concave central cusp than those of OK, which gives these teeth a
shovel-like appearance that is adapted to scooping loose material from between
filamentous algae attached to rocks.

The oral teeth of O. cf. korogwe are much more robust than those of O. 1060 hunteri (Fig. 2.3b; Supplementary Fig. A.2), and they are more closely spaced 1061 and more firmly attached to the jaw bones. All of these observations indicate 1062 that OK is well equipped to scrape filamentous algae from rocks (Yamaoka, 1063 1983). The crowns of its outer-series (bicuspid) teeth are typically obliquely 1064 truncated, with a minor cusp that is taller in comparison to the bicuspids of 1065 OH. This type of tooth crown, and close tooth spacing, gives the outer series 1066 a continuous but notched appearance, like the blade of a jagged knife. The 1067 second series of teeth forms a flattened and even surface of densely spaced 1068 tricuspid teeth, allowing the fish to scrape algae thoroughly from the rocks. 1069 The flattened appearance of second-series tricuspids is due to their specific 1070 crown morphology, with minor cusps that almost equal the major cusp in size. 1071

Both species have a number of unicuspid teeth in the posterior part of the dentary, which probably are new first-generation teeth that were added to the jaw as a result of growth (Huysseune & Witten, 2018). These unicuspid teeth are largely absent from the premaxilla, suggesting that new teeth are first added to the dentary in these *Oreochromis* species.

The functional interpretation of tooth morphology corresponds to the observed patterns of tooth wear. Greater tooth wear is expected in OK than in OH, because of the functional implications of scraping and brushing feeding styles, respectively. Our observation of greater differences in the degree of cusp asymmetry in oral teeth of OK compared to OH is consistent with more substantial wear. The greater bite forces needed for scraping may moreover explain the more robust dentaries of OK (Fig. 2.2).

¹⁰⁸⁴ 2.5.3 Tooth morphology of fossil cichlids from Lake Chala

¹⁰⁸⁵ The high amount of intraspecific morphological variation in the modern-day ¹⁰⁸⁶ Oreochromis species of Lake Chala ideally requires large numbers of fossil teeth to trace morphological change through time (Chapter 6). The limited number of fossil teeth examined in this Chapter therefore urges caution in the interpretation of our results. Nevertheless, the morphospace occupied by the fossil tooth crowns shows a number of aspects that are worth discussing in more detail.

First, fossil teeth recovered from lake sediments of any past time interval 1092 are very likely derived from multiple specimens. Consequently, the broadly 1093 similar morphospace occupation of fossil and modern tooth crowns indicates 1094 that, as anticipated, variation in tooth morphology between individuals of 1095 a species is significantly more limited than the morphological disparity of 1096 teeth within the jaws of a single specimen. This result suggests that teeth 1097 from a limited number of specimens provide a reasonable estimate of the 1098 morphological variation within a population. However, this methodological 1099 convenience comes with a potential disadvantage: if great variation is observed 1100 within populations (groups of interbreeding individuals), but limited variation 1101 exists between populations or species, the opportunities to reconstruct ecological 1102 or evolutionary trends in trophic morphology through time may be limited. 1103 Given that this issue may compromise our primary research objectives, we 1104 explore potential trends in the fossil data further. 1105

Although the morphospace occupied by our collection of fossil teeth does 1106 not fully overlap with those of modern-day O. hunteri in Lake Chala (Fig. 2.8), 1107 the recorded differences between fossil and modern teeth are rather limited 1108 and suggest that the fossil teeth also derive from a large cichlid feeding on 1109 epilithic algae. Our fossil data further revealed a modest but consistent gradual 1110 shift of mean tooth morphology through time, expressed mainly in tricuspid 1111 teeth and towards the morphology of modern-day O. hunteri. In consideration 1112 of the small size and disjunct time line of our current fossil collection, we 1113 refrain from drawing inferences about a possible long-term shift in the feeding 1114 ecology of the *Oreochromis* from Lake Chala. Nevertheless, these results suggest 1115 that the presumed disadvantageous combination of limited variation between 1116 populations with large variation within specimens does not seem to hamper 1117 detection of a long-term trend in tooth morphology, even if the magnitude 1118 of these changes through time is limited in comparison with variation within 1119 assemblages. 1120

1121

The available paleo-environmental (e.g. Verschuren *et al.*, 2009; Barker

et al., 2011) and geochronological (Moernaut et al., 2010; Blaauw et al., 2011) 1122 data indicate that a deep, freshwater aquatic ecosystem with rocky shoreline 1123 existed continuously in the Chala Basin throughout the last 25,000 years; there 1124 is no indication for large ecological disturbances that could have eradicated the 1125 local fish community. Hence, the lack of fossils in the sample from 19,200 year 1126 BP is notable, and more detailed sampling around that period is required to 1127 figure out what may have cause this absence. Although natural colonisation of 1128 an isolated crater lake by fish is considered largely a rare chance event (Elmer 1129 et al., 2012), the presence of cichlids in even very young isolated lakes suggests 1130 that on longtime scales, colonisation of vacant habitat may occur relatively 1131 frequently (Elmer et al., 2010b). In this regard, re-colonisation of Lake Chala 1132 from closely-related parent populations could explain similarity in shape of 1133 fossils before and after a temporary absence. 1134

1135 2.6 Conclusion

This exploratory methodological study demonstrates the potential of fossil 1136 cichlid teeth extracted from lake-sediment cores to trace morphological change 1137 in their trophic apparatus through time. At least in a well-controlled setting 1138 such as Lake Chala, morphological analysis of fossil cichlid teeth may thus 1139 provide a means to elucidate the natural history of local fish communities 1140 including their ecological and evolutionary dynamics. Contingent upon the 1141 availability of greater fossil collections, more intensive and finer-spaced sampling 1142 of fossil teeth will be required to confirm whether the apparent trend in the 1143 morphology of Lake Chala *Oreochromis* is real. Moreover, comprehending the 1144 evolutionary and ecological changes behind the observed morphological trends 1145 will require better understanding of the feeding ecology of modern-day O. 1146 *hunteri* and related species, and its relationship with their trophic morphology. 1147

• The species name *Oreochromis korogwe* is replaced by *O*. cf. *korogwe*, in 1149 accordance with findings in Chapter 3. 1150 • *Tilapia rendalli* is now *Coptodon rendalli* (Dunz & Schliewen, 2013). 1151 • In Material and methods, the rationale behind the selection of the four 1152 fossil assemblages is now addressed (Section 2.3.1). 1153 • The presence of several unicuspid teeth in the posterior part of the 1154 dentaries is now addressed in the Results (Section 2.4.4) and in the 1155 Discussion (Section 2.5.2). 1156 • An intermediate O. cf. korogwe specimen is discussed in more detail in 1157 the Discussion (Section 2.5.2). 1158 • The high number of oral teeth required to overcome intraspecific shape 1159 variation, with reference to Chapter 6, is addressed in the Discussion, 1160 Section 2.5.3. 1161 • Photographs of cleared-and-stained jaws of O. hunteri and O. cf. koroque 1162 have been added to Appendix A. 1163 • Appendix A now contains a table listing the number of teeth per jaw 1164 region included in the morphometric dataset. 1165 • A new figure indicating the location of the different tooth types (unicuspid 1166 - bicuspid - tricuspid) in morphospace is added in Appendix A. 1167

CHAPTER 3

1169 Species integrity and origin of Oreochromis hunteri 1170 1171

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- Modified from**: Dieleman, J., Muschick, M., Nyingi, W.D., & Verschuren, D. (2018). Species 1177
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- Tanzania). Hydrobiologia Online first. doi.org/10.1007/s10750-018-3570-7. 1179
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- 1187
 - ** See Section 3.6 for a list of modifications.

1188 **3.1** Abstract

3

Extensive transfer of tilapia between lakes throughout East Africa has often led 1189 to hybridisation with indigenous fish populations. The endemic Oreochromis 1190 hunteri of Lake Chala, an isolated crater lake near Mount Kilimanjaro, is 1191 potentially susceptible to such introgression from a species formerly identified 1192 as Oreochromis koroque, introduced 30 years ago. We combined whole-body 1193 geometric morphometry on 104 Lake Chala specimens of both taxa with 1194 molecular phylogenetic analysis of mitochondrial loci from 15 O. hunteri and 1195 9 O. cf. koroque specimens to assess whether hybridisation has occurred. 1196 Using fishes from Lake Jipe and Nyumba ya Mungu reservoir, we expanded 1197 our analysis to all four *Oreochromis* species currently inhabiting the Upper 1198 Pangani River system to determine the closest living relative of O. hunteri, and 1199 hence the possible source population of the ancestral species which colonised 1200 Lake Chala. Our results indicate that no interbreeding occurs between O. 1201 hunteri and O. cf. korogwe, and infer O. jipe to be the closest living relative 1202 of O. hunteri. The introduced O. cf. korogwe is a phenotypically uniform 1203 but genetically variable population, the identity of which remains unknown. 1204 The high haplotype diversity of *O. hunteri* is consistent with available fossil 1205 evidence indicating that its ancestor colonised Lake Chala at least 25,000 years 1206 ago. 1207

1208 3.2 Introduction

Tilapia is one of the most productive food fishes in Africa. Tilapia-based 1209 fisheries provide an often indispensable protein source for local food security, 1210 and especially *Oreochromis* species have been utilised for this purpose in global 1211 aquaculture (Eknath & Hulata, 2009). This has led to extensive transfer from 1212 their natural ranges into other regions and countries (Eknath & Hulata, 2009). 1213 The invasive nature of some *Oreochromis* species and their propensity for 1214 hybridisation rightfully has raised concern when they are being introduced 1215 to regions with indigenous tilapine communities (e.g. Agnèse et al., 1998, 1216 D'Amato et al., 2007, Nyingi & Agnèse, 2007, Angienda et al., 2011, Deines 1217 et al., 2014, Ndiwa et al., 2014). It has been suggested that the introgression of 1218 alien genes into local species may also contribute to rapid speciation in cichlids 1219

(Salzburger *et al.*, 2002, Meier *et al.*, 2017b). However, most often it simply
induces a loss of genetic diversity, through the homogenisation of gene pools
(Nyingi & Agnèse, 2007, Crispo *et al.*, 2011, Firmat *et al.*, 2013).

Lake Chala (locally 'Challa', after a nearby village) is a crater lake bridging 1223 the border between Kenya and Tanzania, immediately to the southeast of 1224 Mount Kilimanjaro in East Africa. It harbours the only natural population 1225 of Oreochromis hunteri (Günther, 1889, Seegers et al., 2003), the type species 1226 of the genus Oreochromis Günther, 1889 (Trewavas, 1983). Until the early 1227 1980s it also seems to have been the only fish species inhabiting Lake Chala, 1228 as surveys carried out in 1889, 1902, 1946, 1952 and 1980 did not reveal other 1229 species (Günther, 1889, Dadzie et al., 1988). However, sometime in the late 1230 20^{th} century two other tilapiine species were introduced, namely Coptodon 1231 rendalli Boulenger, 1896 and Oreochromis korogwe Lowe, 1955 (Dadzie et al., 1232 1988, Seegers et al., 2003), as well as a small cichlid identified as Haplochromis 1233 "Chala" (Seegers et al., 2003). The two Oreochromis species have spec. 1234 the potential to hybridise, as the deep open-water environment of the lake 1235 encircled by near-vertical rocky crater walls appears to offer limited potential 1236 for reproductive or niche segregation. 1237

Recent research on the long-term evolutionary and ecological dynamics of 1238 O. hunteri in Lake Chala is based on analysis of its fossil teeth recovered from 1239 the sediment record (Chapter 2, Chapter 6). This study uses the extant O. 1240 *hunteri* population as principal reference for variation in fossil tooth morphology, 1241 and assumes that it is genetically pure. However, introgression of genetic 1242 material due to interbreeding with an introduced species could potentially 1243 impact important morphological features (Parnell et al., 2012, Holzman & 1244 Hulsey, 2017). Such a recently compromised species integrity would complicate 1245 the comparison of modern phenotypes with the fossil record. Quantitative 1246 examination of general body morphology and the shape of oral teeth (Chapter 1247 2) found the two local *Oreochromis* species to be clearly distinct, arguing 1248 against recent and/or ongoing hybridisation. However, demonstration of the 1249 presence or absence of shared haplotypes with molecular-genetic methods 1250 would provide a more sensitive test of hybridisation. Such genetic assessment 1251 is particularly relevant for the endemic O. hunteri population in Lake Chala, 1252 where detection of recent hybridisation would have implications for both 1253 taxonomy and conservation. Also, comparing genetic data of O. hunteri with 1254

that of *Oreochromis* populations from nearby waters might help reveal its phylogenetic associations, by providing information on the possible source population(s) from where the isolated crater lake was colonised, and on the approximate timing of this colonisation (Barluenga *et al.*, 2006, Elmer *et al.*, 2012, Genner & Turner, 2014).

In this Chapter, we aim to (i) validate with genetic evidence the quantitative morphological differences and apparent lack of interbreeding between the endemic and introduced *Oreochromis* in Lake Chala; (ii) identify the closest relative of *Oreochromis hunteri* by molecular phylogenetic analysis of the *Oreochromis* fauna inhabiting Lake Chala and the only two nearby lakes, namely Lake Jipe and the Nyumba ya Mungu reservoir; and (iii) provide an estimate of the timing of *O. hunteri*'s arrival in Lake Chala.

1267 3.2.1 Study area

Lake Chala (3°19'S, 37°42'E) is a freshwater lake with a surface area of 4.5 km² 1268 and a maximum depth that has fluctuated between 92 m and 98 m since 1999 1269 (Wolff *et al.*, 2014). The lake fills a volcanic caldera basin at \sim 880 m above 1270 sea level immediately east of Mount Kilimanjaro, and is hydrologically mainly 1271 influenced by subsurface in- and outflow (Moernaut et al., 2010). Despite 1272 its isolated location, biogeographically it is considered part of the Upper 1273 Pangani River basin (Seegers et al., 2003), which also contains Lake Jipe 1274 (Kenya/Tanzania) and the man-made reservoir Nyumba ya Mungu (Tanzania; 1275 Fig. 3.1). The former is a medium-sized (30 km^2) but very shallow (<3 m) 1276 muddy lake located immediately east of the Pare Mountains. It is fed by 1277 the Lumi River and drains into the Ruvu River, both at the northern end of 1278 the lake, and its shoreline is fringed with swamps. Nyumba va Mungu was 1279 created in 1965, has a surface area of $110 - 180 \text{ km}^2$ depending on rainfall and 1280 drawdown, and gradually increases in depth from 4 m in the north to 41 m 1281 at the dam (Denny, 1978). It lies in the north-south trending valley between 1282 the Lelatema and Pare Mountains, at the former confluence of the Ruvu and 1283 Kikuletwa rivers. The reservoir drains into the Pangani River, which flows into 1284 the Indian Ocean 500 km to the southeast. 1285

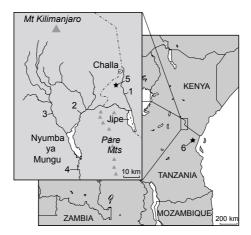


Figure 3.1: Skeleton maps of East Africa and the Upper Pangani River basin (inset) in northern Tanzania, with indication of the sampled surface waters Chala, Jipe and Nyumba ya Mungu. River systems (1-4) and towns (5-6) mentioned in the text are indicated by numbers. 1: Lumi, 2: Ruvu, 3: Kikuletwa, 4: Pangani, 5: Taveta, 6: Korogwe

1286 3.2.2 Regional ichthyofauna of the Upper Pangani system

The indigenous *Oreochromis* fauna of the Upper Pangani region is rather modest. 1287 Apart from O. hunteri, endemic to Lake Chala, O. jipe is considered indigenous 1288 to both Lake Jipe and the Pangani River itself (Lowe, 1955, Trewavas, 1983, 1289 Seegers et al., 2003). Trewavas (1983) also reports the species O. panqani, 1290 with subspecies O. pangani pangani in the Pangani River and O. pangani 1291 girigan in Lake Jipe. However, several recent studies regard O. pangani as 1292 conspecific to O. jipe (Seegers et al., 2003; Seegers, 2008; Froese & Pauly, 2017; 1293 Shechonge et al., 2018), and it is treated here as such. Morphologically, O. 1294 *jipe* is considered the closest relative of O. hunteri, sharing high numbers of 1295 vertebrae, scales and dorsal fin rays (Trewavas, 1983), but to our knowledge 1296 this has never been confirmed by genetic data. Transfer of tilapiine fishes to 1297 improve local fisheries started influencing the ichthyofauna in this region from 1298 the 1970s onwards. For most of these transfers, no written records are available, 1299 and nearly all dates mentioned below are based on observations made during 1300 field surveys, rather than actual accounts. 1301

1302

A survey in 1980 found only the endemic *O. hunteri* in Lake Chala, but

in 1985 catches also included *Coptodon rendalli* and one specimen that the 1303 collectors identified as O. pangani (Dadzie et al., 1988). Since the latter 1304 species does not occur in any recent catches, it either rapidly disappeared 1305 after its introduction, or the specimen was in fact O. hunteri misidentified as 1306 O. panqani. More recently a second Oreochromis species has been found in 1307 significant numbers, and it has been thought to be O. korogwe (Seegers et al., 1308 2003). Morphologically, however, Lake Chala specimens do not fully correspond 1309 to the O. koroque holotype, which naturally inhabits the Lower Pangani River. 1310 Seegers *et al.* (2003) thus recommended to confirm this association, but no 1311 such study has been carried out so far. 1312

The first written account of catches of O. esculentus in Lake Jipe (which 1313 naturally occurs only in Lake Victoria) dates from 1983, followed by C. rendalli 1314 (naturally distributed throughout the Congo River basin, lakes Tanganyika 1315 and Malawi, and southern Africa) in 1985 (Dadzie et al., 1988). However, 1316 Trewayas (1983) suggests that at least O. esculentus must have been introduced 1317 there already in the 1950s, but that it went unnoticed in previous surveys. In 1318 Nyumba ya Mungu, O. esculentus and C. rendalli were caught for the first 1319 time in respectively 1973 and 1974 (Bailey et al., 1978). It is unclear whether 1320 O. *jipe* colonised this reservoir in a natural way, or was introduced (Trewavas, 1321 1983). 1322

This general lack of written records complicates determining the exact 1323 source populations of introduced species, but some inferences can be made. 1324 Probably, most of the region's lakes were stocked with fish from nearby fish 1325 ponds, from which transporting living fry is more probable and logistically 1326 more feasible. The fish ponds still present today around the town of Taveta and 1327 near Lake Jipe are good candidates, as they date back to the late 1940s when 1328 the British colonel Ewart Scott Grogan settled in the region and established a 1329 sisal farm that included such ponds (Dadzie et al., 1988). Tilapiines were also 1330 reared in governmental fish ponds at the town of Korogwe, located downstream 1331 along the Pangani River, for stocking reservoirs and ponds throughout the 1332 Tanganyika region (Lowe, 1955, Bailey et al., 1978). Among those species were 1333 O. jipe and O. koroque, but also O. esculentus and O. variabilis from Lake 1334 Victoria, and C. rendalli and O. macrochir that had previously been raised in 1335 ponds in the D.R. Congo province of Katanga (Lowe, 1955, Dadzie *et al.*, 1988). 1336 It seems thus most plausible that species introduced in the Upper Pangani 1337

Region, including Lake Chala, derive from populations that were reared in theKorogwe or Taveta ponds.

$_{1340}$ 3.3 Methods

1341 3.3.1 Specimen collection

For this study, we obtained 104 specimens from local fishermen servicing nets 1342 on the Kenya (southeast) side of Lake Chala, in September 2012, January 2014 1343 and September 2015. Similarly, 15 specimens were obtained from fishermen 1344 along the southeast shore of Lake Jipe and 10 from the north shore of Nyumba 1345 ya Mungu (hereafter, NyM) during the same periods. Initially the local names 1346 given to the diverse species were recorded for each specimen; the corresponding 1347 nominal species names were assigned afterwards. From Lake Chala, two 1348 Oreochromis species were distinguished, being O. hunteri ('Chala') and O. cf. 1349 koroque ('Bandia'). From Lake Jipe, we collected O. jipe ('Asilia') and O. 1350 esculentus ('Polana'). At NyM we collected O. jipe ('Asilia') and O. esculentus 1351 ('Polana') (Supplementary Table B.1). 1352

1353 3.3.2 Geometric-morphometric analysis

Whole-body photographs of fresh specimens were taken in the field prior to preservation, from a perpendicular angle. This was carried out by positioning specimens on graph paper with their left side facing up and fins spread out.

Overall body morphology was analysed as in Chapter 2, using 16 tradi-1357 tional landmarks (Fig. 2.1) digitised in tpsDig2 version 2.17 (Rohlf, 2013a). 1358 The present chapter includes data on 95 Lake Chala specimens previously 1359 presented in Chapter 2. Size was calibrated using the graph paper visible on 1360 the photograph. The digitised landmark dataset was aligned via Procrustes 1361 superimposition in the program CoordGen6h of the Integrated Morphometrics 1362 Package software (Sheets, 2008). Procrustes coordinates and rescaled centroid 1363 size were saved as data matrix file in IMP format. All further analyses were 1364 performed on these Procrustes coordinates. 1365

Overall variation in body shape was analysed using principal component analysis (PCA) in R (R Development Core Team, 2016), and the number of principal components to consider was assessed using scree-plot analysis and a 3

broken-stick model. Multivariate Analysis of Variance (MANOVA) with four 1369 constraints was used to determine whether the fish taxa as distinguished by their 1370 local names indeed reflect significantly distinct shape morphs, and Canonical 1371 Variate Analysis (CVA) was used to determine the axes of maximal group 1372 separation. Both analyses were conducted on minimal Mahalanobis distances 1373 in PAST v.2.17 (Hammer et al., 2001), and results were cross-validated by 1374 comparing leave-one-out (jack-knifed) classifier tables to the original confusion 1375 matrix. Substantial differences between these two classifier tables would reflect 1376 the important influence of one or few specimens on the observed outcome, and 1377 hence reveal any unreliable results. 1378

1379 3.3.3 MtDNA genotyping & analysis

We isolated DNA from tissue sample collected and stored in the field in 1380 99.5% absolute ethanol using the Blood and Tissue DNA isolation kit (Qiagen) 1381 following the manufacturer's specifications. Two widely-used mitochondrial loci 1382 were chosen for this study, NADH dehydrogenase subunit 2 (ND2) and a fast-1383 evolving part of the control region (CR), allowing straightforward comparison 1384 with GenBank sequences available for other *Oreochromis* species. These loci 1385 were amplified via polymerase chain reaction (PCR), using the published 1386 primers MET and TRP for ND2 (Kocher et al., 1995), and L-Pro-F (Meyer 1387 et al., 1994) and TDK-D (Lee et al., 1995) for CR. PCR was performed in 25 μ L 1388 reaction volumes, containing 1 μ L DNA extract, 2.5 μ L PCR buffer II (Applied 1389 Biosystems), 0.5 μ L of either primer [10 μ M], 1 μ L MgCl₂ solution [25 mM], 0.4 1390 μL dNTP solution [10 mM], 0.1 μL AmpliTag Gold DNA polymerase [5U/ μL] 1391 (Applied Biosystems), and 19 μ l water. Thermocycling was performed with an 1392 initial denaturation for 3 minutes at 95°C, then 30 cycles with 30 seconds at 1393 95°C, 30 seconds at 55°C, and 1 minute at 72°C, followed by final elongation 1394 for 7 minutes at 72°C. Reaction products were cleaned up using ExoSAP-IT 1395 PCR Product Cleanup Reagent (Affymetrix), following the manufacturer's 1396 instructions. Purified PCR products were used for cycle sequencing reactions 1397 using the BigDye Terminator Mix v3.1 (Applied Biosystems). Cycle sequencing 1398 was performed for each sample and primer combination using 4 μ L of BigDye 1399 Terminator 3.1 Ready Reaction Mix, 1.5 μ L of primer [10 μ M], 2 μ L cleaned 1400 PCR product, and 2.5 μ L water. Thermocycling was performed with 1 minute 1401 at 96°C, then 25 cycles with 10 seconds at 94°C, 5 seconds at 50°C, and 4 minutes 1402

at 60°C. Cycle sequencing reactions were cleaned using ethanol precipitation.
Sanger sequencing was performed on an ABI 3730 48-well capillary DNA
Analyser (Applied Biosystems, Foster City, CA, USA). Electropherograms and
their automatic translation were checked by eye and trimmed. Overlapping
sequence reads from either direction were merged for each sample and locus.
For each locus, 38 sequences were generated for this study (Supplementary
Table B.1).

In total, four sequence alignments were generated. Lengths differed be-1410 tween alignments due to the occurrence of gaps with more distantly related 1411 species, and as a result of trimming positions with excess missing data from the 1412 alignment ends. O. hunteri CR sequences (N=15) were aligned (430 positions) 1413 and used for demographic analyses. All new Oreochromis sp. sequences (N=38) 1414 were aligned for each locus separately (1051 positions for ND2, 435 for CR), and 1415 concatenated to generate a haplotype network. Additional, published sequences 1416 of each locus were downloaded from Genbank (https://www.ncbi.nlm.nih.gov, 1417 Supplementary Tables B.2 and B.3) and analysed jointly with those new to this 1418 study. The datasets were pruned and trimmed, resulting in two alignments with 1419 1040 and 437 positions, and 71 and 90 sequences, for ND2 and CR, respectively. 1420 These alignments were used for locus-wise phylogenetic analysis in RAxML 1421 (v.8.2.4; Stamatakis, 2014). For each locus, a separate maximum-likelihood 1422 tree was generated to make best use of available GenBank sequences. To 1423 find the best-scoring tree in each case, we performed rapid bootstrap analyses 1424 using the GTRGAMMAI model of sequence evolution and 100 alternative 1425 runs from distinct starting trees. The resulting trees were visualised in, and 1426 figures created with, FigTree v.1.4.3 (Rambaut, 2009) and fitchi (Matschiner, 1427 2016), for the locus-wise molecular phylogenies and the haplotype network. 1428 respectively. 1429

To test for a genetic signature of past population expansion in O. hunteri's 1430 CR sequences, we performed a haplotype mismatch distribution analysis in 1431 Arlequin v.3.5 (Excoffier et al., 2005) and a coalescent Bayesian skyline plot 1432 analysis (BSP; Drummond et al., 2005) in BEAST2, v.2.4.5 (Bouckaert et al., 1433 2014). Populations that have undergone a period of sudden or exponential 1434 growth in the past exhibit a characteristically unimodal, wave-like pattern in the 1435 distribution of haplotype mismatches. If present, the mode of this distribution 1436 together with estimates of generation time and mutation rate can be used to 1437

3

infer the approximate timing of population expansion (e.g., Barluenga et al., 1438 2006). We performed the BSP analysis using a strict molecular-clock model 1439 with the base-substitution-rate estimate of 0.0324 changes per site per million 1440 years (SE 0.0139) of Genner et al. (2010). Our analysis employed a non-coding 1441 site model, and the Hasegawa-Kishino-Yano (HKY) substitution model with 1442 empirical base frequencies (Hasegawa et al., 1985), identified as best choice for 1443 these data, using jModelTest v.2.1.10 (Darriba et al., 2012). Chain length was 1444 25,000,000 steps, and the first 10% were discarded as burn-in. This analysis 1445 also estimated the timing of the deepest coalescence event of the O. hunteri 1446 CR sequences. 1447

1448 **3.4** Results

¹⁴⁴⁹ 3.4.1 Morphological characterisation of regional tilapiine taxa

Although the broken-stick model indicates that 6 principal components (PCs) 1450 are significant, the scree plot (Supplementary Figure B.1) shows that only 1451 PC1-3 explain at least 10% of the morphological variation. Therefore, we only 1452 retained principal components explaining at least 10% of the morphological 1453 variatio (PC1-3). The first two axes of the PCA (Fig. 3.2a) together explain 1454 47.14% of the observed variation in general body morphology. Principal 1455 component 1 (PC1) mainly reflects variation in body elongation and depth. 1456 with long and slender bodies on the positive side and shorter, deeper bodies 1457 on the negative side of the axis (Fig. 3.2c). PC1 distinguishes O. hunteri 1458 from the other *Oreochromis* species. PC2 reflects differences within body 1459 depth: specimens with positive PC2-values are ventrally flattened, whereas 1460 negative PC2-values correspond to ventrally-extended specimens (Fig. 3.2d). 1461 PC3, which explains an additional 12% of the observed variation, captures 1462 differences in the length of the head: specimens with positive PC3-values have 1463 shorter heads (i.e., relatively smaller distance between mouth and operculum). 1464 whereas specimens with negative PC3-values have more elongated heads. 1465

Although assumptions for CVA and MANOVA were not met, jack-knifed confusion matrices did not differ substantially from the original classification table. Therefore, we consider these results as reliable, except for *O. esculentus*, where the small sample sizes (2 specimens from Jipe, 4 from NyM) prevented comparison with the other species except for *O. hunteri*. Uncorrected and Bonferroni-corrected pairwise comparison results indicate that *O. hunteri*, *O.* cf. *korogwe* and *O. jipe* specimens as identified by their local names differ significantly in overall body morphology at the (overall) 5% confidence level (Table 3.1; Wilks' $\lambda = 0.003$, p < 0.001); the CVA scatter plot (Fig. 3.2b) shows all four taxa to be clearly distinct from one another. *O. hunteri* is separated from the other species along the first axis (CV1), whereas the other species are separated along the second axis (CV2).

Table 3.1: Uncorrected (above the diagonal) and Bonferroni-corrected (below the diagonal) *post-hoc* results, showing that all taxa as identified by their local names differ significantly in overall body morphology. The limited sample size of Polana caused pairwise comparisons to be non-applicable (N.A.) in some cases

		'Chala'	'Asilia'	'Bandia'	'Polana'
O. hunteri	'Chala'		< 0.001	< 0.001	< 0.001
O. jipe	'Asilia'	< 0.001		< 0.001	N.A.
O. cf. korogwe	'Bandia'	$<\!0.001$	0.003		N.A.
O. esculentus	'Polana'	$<\!0.001$	N.A.	N.A.	

¹⁴⁷⁸ 3.4.2 Genetic characterisation of regional *Oreochromis* taxa

1479 Phylogenetic trees

Figures 3.3 and 3.4 depict the phylogenetic trees of a selection of relevant 1480 African *Oreochromis* taxa based on sequence data from the mitochondrial 1481 control region (CR, Fig. 3.3) and NADH dehydrogenase subunit 2 (ND2, Fig. 1482 3.4) loci, respectively. Specimens of the four *Oreochromis* taxa as identified by 1483 local fishermen, and as validated by geometric-morphometric analysis, largely 1484 cluster into distinct mitochondrial clades. Further, CR and ND2 sequences 1485 suggest very similar phylogenetic affiliations, as could be expected from a 1486 non-recombining pair of markers. Only few of the studied specimens appear 1487 genetically distinct from these clusters. Specimens representing 'Chala' (O. 1488 hunteri) from Lake Chala and 'Asilia' (O. jipe) from both Lake Jipe and NyM 1489 cluster together in one clade, but with a distinct, albeit nested, split between 1490 the two species. Comparison with relevant GenBank sequences confirms high 1491 sequence similarity of 'Asilia' to gene fragments of O. jipe and suggests a close 1492 relationship with O. amphimelas specimens from an unknown locality. The 1493 latter species occurs naturally in lakes of the Eastern Rift Valley in Tanzania 1494

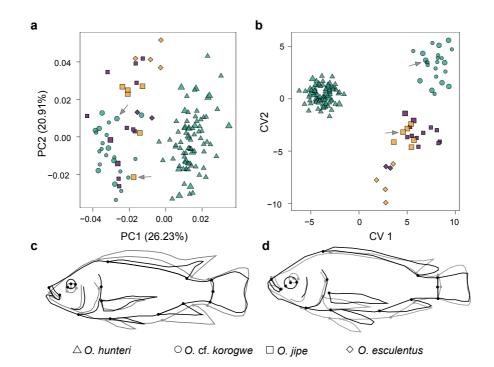


Figure 3.2: a PCA ordination plot synthesising overall variation in body morphology among the four *Oreochromis* species currently inhabiting the Upper Pangani River basin. Green, purple and yellow symbols represent specimens from respectively Lake Chala, Lake Jipe and NyM reservoir, with large symbols of each type and colour indicating sequenced specimens. The two arrows point to the specimens of *O. jipe* and *O.* cf. *korogwe* with genetically distinct positions in the phylogenetic trees. **b** CVA scatterplot showing maximal phenotypic separation of the four groups of specimens attributed to each of the four taxa. The outline drawings represent the specimens at the **c** positive (black) and negative (grey) extremes of PC1, and **d** the positive (black) and negative (grey) extremes of PC2.

(Manyara, Eyasi, Kitangiri, Singida), situated to the west of the Upper Pangani
region. However, CR-sequences of *O. amphimelas* from Lake Manyara (clade
containing *O. amphimelas* AF296490) occur in very different parts of the tree.
Nevertheless, the latter clade is separated from *O. hunteri* and *O. jipe* by
nodes with very low support values, so that their close relationship with *O. amphimelas* cannot be ruled out.

One 'Asilia' specimen from NyM (ASILIA nym011) clusters with *O. niloticus* genotypes, although morphologically it groups with the other 'Asilia' examined in this study (Fig. 3.2a). The Lake Chala fishes identified by local

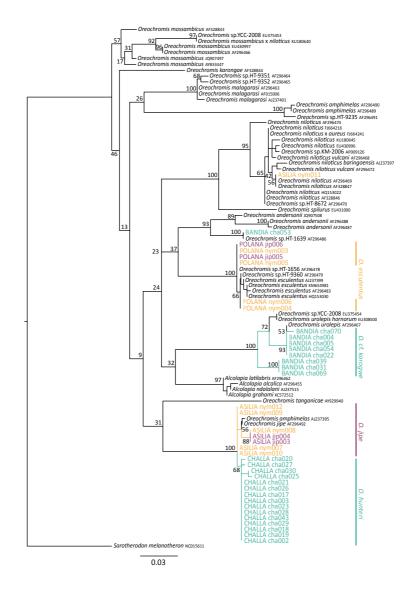


Figure 3.3: Maximum-likelihood phylogenetic tree of relevant African *Oreochromis* taxa based on sequences of the control region (CR) locus in mtDNA. *Sarotherodon melanotheron* was used as outgroup. The specimens sequenced in this study are colour-coded per site, as in Fig. 3.2: Chala (green), Jipe (purple) and NyM (yellow). Node support values are given in percent and are based on 100 bootstrap replicates. Note the low support for deeper nodes, but high support for those nodes on which identification of the closest relatives of the target species relies

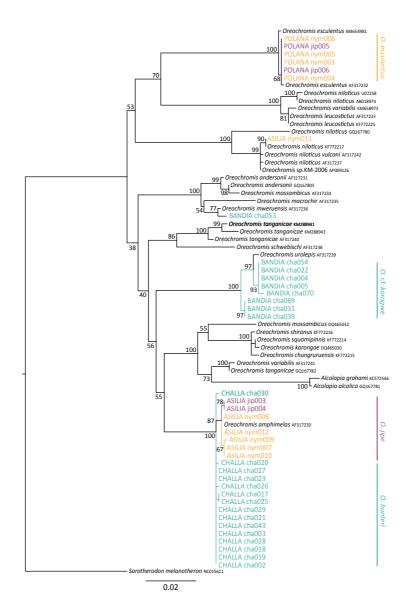


Figure 3.4: Maximum-likelihood phylogenetic tree based on sequences of the NADH dehydrogenase subunit 2 (ND2) locus in mtDNA. *Sarotherodon melanotheron* was used as outgroup. The specimens sequenced in this study are colour-coded per site, as in Fig. 3.2: Chala (green), Jipe (purple) and NyM (yellow). Node support values are given in percent and are based on 100 bootstrap replicates. Note the low support for deeper nodes, but high support for those nodes on which identification of the closest relatives of the target species relies

fishermen as 'Bandia' (O. cf. koroque), although morphologically uniform 1504 (Fig. 3.2a), display a distinct split within their main clade, with CR and ND2 1505 sequences separating the same two sub-groups of specimens. This differentiation 1506 appears greater than the variation within O. hunteri, or even the difference 1507 between O. hunteri and O. jipe (Fig. 3.3 & 3.4). Most of our 'Bandia' 1508 specimens cluster phylogenetically with O. urolepis. One available O. urolepis 1509 CR sequence, representing a specimen from the Wami river, was even identical 1510 to a subset of 'Bandia' sequences. This river belongs to the natural range 1511 of O. urolepis, and is located just south of the Pangani River basin. One 1512 particular Lake Chala specimen (BANDIA cha053) is phylogenetically even 1513 further removed from the main 'Bandia' clade. Our phylogenetic reconstruction 1514 based on the CR gene suggests close affinity with an unidentified Oreochromis 1515 specimen (HT-1639) collected from Pangani River, and dubbed Oreochromis 1516 'Korogwe' in the original publication (Nagl et al., 2001). Its ND2-sequence 1517 reveals that a specimen of O. mweruensis, collected in Lake Mweru-Wantipa 1518 (Zambia), may be a close relative (Klett & Meyer, 2002). 1519

1520 Haplotype network

An unrooted haplotype network of all genotyped *Oreochromis* specimens based 1521 on concatenated CR and ND2 sequences reveals 20 different haplotypes (Fig. 1522 3.5). Again, the four taxa locally recognised as such (and distinguished by 1523 morphology, Fig. 3.2a) also group into distinct mitochondrial clusters. Among 1524 our 15 O. hunteri specimens, haplotype 20 is the most abundant (9 specimens), 1525 with six less common haplotypes (one specimen each) differing from haplotype 1526 20 by up to five mutation steps. These seven O. hunteri haplotypes share 1527 a hypothetical common ancestor (haplotype 22) with O. jipe. Our seven O. 1528 *jipe* specimens comprise five haplotypes, of which one is found in Lake Jipe 1529 (2 specimens) and four in NyM, three of which are more closely related to O. 1530 hunteri (Fig. 3.5). 1531

The deep split in the main group of Lake Chala 'Bandia' (*O. cf. korogwe*) specimens separates haplotype 16 (3 specimens) from haplotypes 14, 15 and 17 (together 5 specimens), and reiterates the observations in the phylogenetic trees. Haplotypes of the aberrant 'Bandia' specimen from Lake Chala (12; near *O. mweruensis*) and 'Asilia' specimen from NyM (10; near *O. niloticus*) are also here strongly isolated from their respective, morphology-based, clusters.

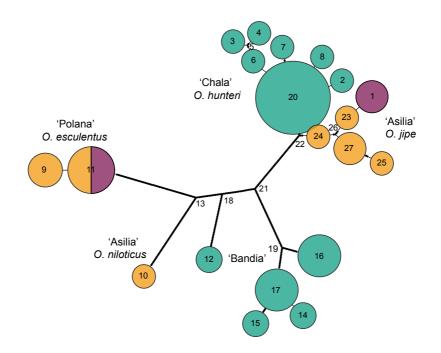


Figure 3.5: Unrooted haplotype network of all 38 *Oreochromis* specimens genotyped in this study, based on mitochondrial CR and ND2 sequences. Each circle represents one haplotype, numbered 1 to 27 including the seven hypothetical haplotypes at network nodes. The diameter of the circles is proportional to the number of individuals with a certain haplotype (1 to 9); the black dots represent undocumented mutation steps between the haplotypes. Colour codes are by water body, as in Figs. 3.2, 3.3 and 3.4: Chala (green), Jipe (purple) and NyM (yellow)

¹⁵³⁸ Mismatch distribution and skyline plot

The observed haplotype mismatch distribution of the 15 genotyped O. hunteri 1539 specimens is not unimodal (Fig. 3.6a), hence no distinct event of past popula-1540 tion expansion can be defined. Likewise, the Bayesian skyline plot (Fig. 3.6b) 1541 does not provide evidence of a sudden large change in population size that 1542 might indicate a post-colonisation expansion. Instead, both results suggest 1543 that the size of the Lake Chala population has been relatively stable, or only 1544 slightly and continuously increasing, over an extended period of time. We 1545 estimate the oldest coalescence event for the CR sequences to have occurred 1546 approximately 100,000 years ago (100 ka), with a median of 92.4 ka and a 95%1547 highest posterior density (HPD) interval ranging from 230 ka to 19 ka. 1548

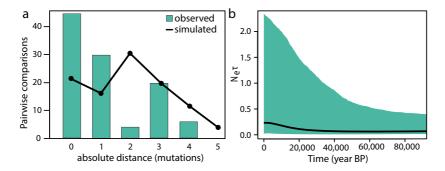


Figure 3.6: a Empirical distribution of haplotype mismatches among the CR sequences of 15 *O. hunteri* individuals (green bars) and their expected distribution from simulations (black line). b Bayesian Skyline Plot visualising the modelled median effective female population size (NeT, scaled by generation time τ) versus time in years before present (BP)

1549 3.5 Discussion

1550 3.5.1 Oreochromis in Lake Chala

To allow treatment of the extant O. hunteri as a modern-day reference for 1551 the fossil record of Lake Chala, hybridisation with species recently introduced 1552 to Lake Chala should ideally be ruled out. Although clear morphological 1553 distinction between O. hunteri and O. cf. korogwe was demonstrated in Chapter 1554 2 and is confirmed by this study (Fig. 3.2), a complementary molecular-genetic 1555 approach enabled direct assessment of the likelihood of past hybridisation events. 1556 The results of our phylogenetic analyses (Figs. 3.3, 3.4 & 3.5) show that Lake 1557 Chala specimens assigned to these two *Oreochromis* species form distinct, well-1558 supported genetic clades and do not share mitochondrial haplotypes, suggesting 1559 that hybridisation in Lake Chala is absent or at least very rare. Hybridisation 1560 can be detected more reliably with a combination of mitochondrial and nuclear 1561 markers, as introgression does not necessarily always affect both genomes at 1562 the same time (Nyingi & Agnèse, 2007, Angienda et al., 2011). However, 1563 mitochondrial introgression appears to happen more readily than introgression 1564 of nuclear loci, due to, for example, interspecific matings being more likely to 1565 occur between females of the (initially) rare invading species and males of the 1566 normally more abundant native species (Wirtz, 1999). Mitochondrial alleles 1567 might also be comparatively neutral in a new genetic background, as opposed 1568

to alleles of nuclear genes, or of loci those genes are linked to (Martinsen 1569 et al., 2001). In Oreochromis, cases of mitochondrial introgression have been 1570 demonstrated to occur without apparent nuclear introgression (Rognon & 1571 Guyomard, 2003, Nyingi & Agnèse, 2007), or with parallel introgression of only 1572 a few nuclear loci (Ndiwa et al., 2014). While this discrepancy might in part 1573 reflect a bias towards traditionally-studied mitochondrial loci, these studies 1574 at least demonstrate that mitochondrial introgression in *Oreochromis* appears 1575 to occur readily and frequently. As O. hunteri and O. cf. koroque from Lake 1576 Chala can easily be distinguished phenotypically and are not found to share 1577 mitochondrial haplotypes despite three decades of syntopy, we consider the 1578 species integrity of O. hunteri to be currently intact, and regard its morphology 1579 as a reliable modern-day reference frame for interpretation of its fossil record. 1580

The uniform and spatially contiguous habitat of Lake Chala does not 1581 provide much opportunity for species segregation, but the observed lack of 1582 mitochondrial exchange between the native and the introduced Oreochromis 1583 does suggest some form of reproductive isolation. Postzygotic barriers are 1584 not very prominent in closely-related cichlid species (Stelkens et al., 2010), 1585 but various forms of premating isolation may prevent the two species from 1586 interbreeding. Visual identification of species-specific coloration patterns, 1587 olfactory cues and sound recognition are important segregation mechanisms 1588 in cichlid species (Fryer & Iles, 1972). In fact, Stelkens & Seehausen (2009) 1589 found that phenotypic divergence predicts assortative mating better than does 1590 genetic distance between species. Visual cues may be less important in tilapias. 1501 but other mechanisms of isolation have been suggested in this group, such as 1592 separation in spawning time and distinctions in microhabitat preference (Pullin 1593 & Lowe-McConnell, 1982, Lowe-McConnell, 1987, Beveridge & McAndrew, 1594 2000). 1595

Although distinction between species within tilapiine genera is often 1596 notoriously difficult (Nagl et al., 2001), the phenotypic/genotypic clusters of 1597 the four species analysed in this study coincide largely with usage of their local 1598 names, indicating that fishermen in each lake most often differentiate accurately 1599 and consistently between these fish taxa. However, one 'Asilia' specimen (in 1600 principle O. *jipe*) has a mitochondrial genotype clustering with O. *niloticus*, 1601 despite its phenotypical similarity to O. jipe. Our morphometric dataset does 1602 not contain O. niloticus specimens, and although this species has never been 1603

encountered in past Lake Jipe surveys (Dadzie et al., 1988, Seegers et al., 2003), 1604 the possibility exists that it has been introduced in recent years, and that this 1605 specimen is an actual O. niloticus or an individual carrying an introgressed 1606 mitochondrial haplotype. Importantly, the phenotypically uniform 'Bandia' 1607 (O. cf. koroqwe) in Lake Chala also contains one individual with a very distinct 1608 genotype and a pronounced subdivision of the other specimens into two clades. 1609 A possible explanation for the occurrence of phenotype-genotype mismatches 1610 is that the aberrant mitochondrial genotype introgressed via an interbreeding 1611 event, either ancient or recent (Rognon & Guyomard, 2003, Ndiwa et al., 1612 2014). This process does not explain the genotypic division within the main 1613 'Bandia' cluster, however. Historical collections suggest that O. cf. koroque 1614 has been introduced to Lake Chala only during the 1980s (Dadzie *et al.*, 1988). 1615 Therefore, it is unlikely that this deep split, more pronounced than even the 1616 split between O. hunteri and O. jipe, has arisen locally in such a short time, 1617 while the morphological uniformity of 'Bandia' implies that the population 1618 indeed consists of one species. We propose that multiple stocking events, or 1619 one single event containing a mix of genotypes, are a likely cause for these 1620 distinct genotypes currently coexisting in Lake Chala. 1621

There also remains doubt about the true identity of 'Bandia', which Seegers 1622 et al. (2003) attributed to O. korogwe with some reticence. Whereas all 9 1623 specimens genotyped in this study are morphologically similar (Fig. 3.2a), most 1624 of their CR- and ND2-sequences cluster with O. urolepis. The one exception 1625 has a CR-sequence most similar to that of a specimen identified by Nagl et al. 1626 (2001) as O. 'Korogwe', but this only refers to the eastern Tanzanian village 1627 where it was collected (in Genbank this specimen is listed as *Oreochromis* sp.). 1628 The ND2-sequence of the same specimen suggests close relationship with a 1629 completely different species, O. mweruensis. No other O. koroqwe sequences 1630 are currently available, and at this point we cannot rule out that our sequences 1631 would align with other O. koroque specimens. Our data nevertheless suggest 1632 that the Lake Chala 'Bandia' (O. cf. korogwe in this paper) were stocked from 1633 at least two fish ponds, each containing a distinct O. urolepis or O. koroque 1634 population which itself may already have undergone prior interbreeding with 1635 other Oreochromis taxa. Future genetic studies including nuclear markers 1636 could elucidate the exact identity of 'Bandia'. 1637

¹⁶³⁸ 3.5.2 The likely ancestor of *O. hunteri*

Trewavas (1983) first suggested that O. jipe and O. hunteri might be closely 1639 related, on the basis of their similar number of vertebrae, which is generally 1640 higher (31-34) than in other Oreochromis species such as O. esculentus (30-31). 1641 One would therefore expect the Upper Pangani tilapiines to have relatively 1642 elongated bodies, and to cluster together in a PCA where body elongation is an 1643 important character separating phenotypes along the principal axis of variation 1644 (PC1). Our geometric-morphometric data do not support this suggestion. 1645 Based on morphological data alone (Fig. 3.2), O. hunteri is distinct from 1646 the other three species, whereas O. jipe clusters with O. cf. korogwe and O. 1647 esculentus. 1648

Our molecular-phylogenetic analyses, in contrast, do reveal O. *jipe* and 1649 O. hunteri to be each other's closest relative (Figs. 3.3 & 3.4). The two 1650 species form a polytomy in both phylogenetic trees, but based on CR-sequences 1651 O. hunteri is nested within O. jipe, whereas the opposite is true for ND2. 1652 This indicates that the employed markers may not be optimal to resolve the 1653 exact relationship of the two species, and we wish to be prudent with further 1654 interpretations. Nevertheless, genetic diversity of the O. jipe indigenous to Lake 1655 Jipe seems to be nested in the greater genetic diversity of modern-day O. jipe 1656 from NvM for both markers, despite the fact that the lacustrine habitat of NvM 1657 reservoir is only 50 years old. We cautiously suggest that this high diversity 1658 reflects standing variation retained through time in riverine populations of O. 1659 *jipe*, which seeded the newly formed NyM as well as the natural but climate-1660 sensitive Lake Jipe, after a (relatively recent) environmental perturbation had 1661 eradicated its lacustrine population. In this context, we follow Seegers et al. 1662 (2003) in considering the riverine O. pangani as conspecific with the lacustrine 1663 O. *jipe*, notwithstanding some morphological differences in oral and pharyngeal 1664 teeth that had led Bailey et al. (1978) and Trewavas (1983) to describe them 1665 as two distinct species. Although no O. pangani specimens were available for 1666 us to address this issue with genetic analyses, we surmise that O. panqani 1667 may well be the riverine representative of O. jipe, which ensured the species' 1668 survival in the Upper Pangani River basin through past episodes of climatic 1669 drought when the regions shallow lakes fell dry. Although O. *jipe* clusters 1670 with an O. amphimelas specimen in both trees, most well-described specimens 1671 of the latter occur in other parts of the tree, and that particular specimen 1672

may be misidentified or may have been subject of mitochondrial introgression.
Therefore, also the apparent relationship between *O. amphimelas* and *O. jipe*should be treated with caution.

¹⁶⁷⁶ 3.5.3 Timing and mode of the colonisation of Lake Chala

The colonisation of isolated crater lakes by fish is still an enigmatic process 1677 (Barluenga & Meyer, 2010, Elmer et al., 2012). The main mechanisms consid-1678 ered are human introduction, a hypothetical former aquatic connection and 1679 natural introduction by air (Elmer et al., 2012). Although undocumented, 1680 late- 20^{th} century human introduction is almost certainly how O. cf. korogwe 1681 and C. rendalli arrived in Lake Chala. Yet fossil fish teeth and bones occurring 1682 throughout the presently recovered part of the sediment record reveal presence 1683 of O. hunteri in Lake Chala since at least 25,000 years ago (Chapter 6). Ancient, 1684 or at least pre-colonial, stocking of fishless lakes in this region of East Africa, 1685 if it did occur, was most likely restricted to the period after ca. 1000 AD, 1686 when Bantu farmers first settled in the Mt. Kilimanjaro region (Håkansson, 1687 2008). The high-rimmed volcanic caldera basin holding Lake Chala was formed 1688 after roof collapse of a magma chamber of Mt. Kilimanjaro, and hence it has 1689 always been hydrographically isolated from Upper Pangani tributary streams. 1690 Therefore, the introduction of fish by air, such as the transfer of fertilised eggs 1691 by birds, arguably remains the only plausible explanation of how the ancestor 1692 of *O. hunteri* arrived in Lake Chala. 1693

Assuming that the niche space available to the colonising O. hunteri 1694 ancestor was not filled by other (now locally extinct) fish species, the ancestral 1695 O. hunteri population probably expanded rapidly after this initial colonisation. 1696 If so, the genetic signature of this ancient population expansion in today's O. 1697 hunteri population should provide an estimate of the time passed since then. 1698 However, as neither the haplotype mismatch distribution nor the Bayesian 1699 skyline plot of the 15 O. hunteri specimens we sequenced (Fig. 3.6) reveal 1700 an unambiguous signal of rapid population expansion, we cannot conclude 1701 with certainty that such rapid population expansion has actually occurred. 1702 Although sample sizes similar to ours have allowed the detection of past 1703 population expansions in some studies (e.g., Genner & Turner, 2014), some 1704 authors suggest that sample sizes must be on the order of 20-40 (Drummond 1705 & Bouckaert, 2015) or even 50 (Grant, 2015) for this purpose. Estimating the 1706

approximate timing of the putative population expansion which followed the
colonisation of Lake Chala by the ancestor of *O. hunteri*, using a coalescence
approach, must hence await the sequencing of additional specimens.

The structure of the modern-day haplotype network of O. hunteri, in 1710 which 15 specimens yield seven haplotypes with up to five mutations between 1711 them, does suggest that this endemic population is relatively ancient, i.e., 1712 in line with the fossil evidence. The age of Lake Chala itself is estimated at 1713 approximately 250,000 years, based on the total depth of its sedimentary record 1714 as revealed by seismic-reflection stratigraphy (Moernaut et al., 2010) relative 1715 to the radiocarbon-dated upper portion of this record (Verschuren et al., 2009, 1716 Blaauw et al., 2011). The HPD interval derived from our genetic data suggests 1717 the age of the oldest coalescence event within O. hunteri to range between 1718 230,000 and 19,000 years. Although this is a rather wide interval of time, its 1719 upper (older) end is consistent with the current best estimate of the age of 1720 Lake Chala, whereas its lower (younger) end is only a slight underestimation of 1721 the minimum age of the population based on fossil evidence. Given the modest 1722 number of sequences currently available, and therefore the potential for more 1723 distant haplotypes to remain undocumented at this time, we consider an early 1724 rather than late colonisation most plausible. 1725

1726 3.6 List of modifications

1727 1728	• The conspecific nature of <i>O. jipe</i> and <i>O. pangani</i> is discussed in more detail in the Introduction, Section 3.2.2.
1729 1730 1731	• An objective method to determine the number of principal components (PCs) to consider is added to Material and methods, Section 3.3.2. The results are discussed in Section 3.4.1, and visualised in Appendix B.
1732 1733	• An explanation for the choice of mitochondrial loci is added to Material and methods, Section 3.3.3.
1734 1735	• The reasoning behind the generation of separate trees for the two loci is added to Material and methods, Section 3.3.3.
1736 1737	• The caption of Fig. 3.2 is changed to explain how the outline drawings were generated.
1738 1739	• A more detailed explanation of the origin of Lake Chala has been added to the Discussion, Section 3.5.3.

CHAPTER 4

1741	
1742	Feeding habits of endemic Oreochromis hunteri and
1743	introduced tilapiines in Lake Chala
1744	

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- 1749

1740

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of Fish Biology

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1756 4.1 Abstract

Two tilapine cichlid species that were recently introduced into isolated crater 1757 lake Chala (Kenya/Tanzania) may interfere with the endemic and critically 1758 endangered Chala tilapia (Oreochromis hunteri Gnther, 1889) by increased 1759 competition for food. However, nothing is known about the feeding habits of 1760 these fish species in Lake Chala, hampering a comprehensive assessment of the 1761 threat these introduced species may pose. Here, we describe the feeding habits 1762 of both the endemic O. hunteri and introduced Oreochromis cf. korogwe and 1763 Redbreast tilapia (Coptodon rendalli (Boulenger, 1897)), with special focus 1764 on seasonal changes in food availability and interspecific differences. We used 1765 a combination of gut content analysis (n=36) and stable-isotope composition 1766 (n=17) to asses feeding behaviour over 20 consecutive months. We compared 1767 these samples to the limnetic phytoplankton composition in these months, and 1768 to the composition of littoral food sources. We found that all species feed in the 1769 littoral zone for most of the year, but migrate to open water in response to an 1770 annual phytoplankton bloom. However, interspecific differences were apparent 1771 in fish stable-isotope composition, and potentially reflect differences in food 1772 preference during the phytoplankton bloom. We suggest that O. hunteri is 1773 making better use of feeding opportunities provided by the lake, and may thus 1774 retain its competitive advantage over recent newcomers. 1775

1776 4.2 Introduction

Anthropogenic introductions of non-native species into natural systems are a 1777 major concern in conservation biology, as they often bring considerable hazards 1778 for the system's indigenous populations (Vitousek et al., 1997; Strayer, 2010). 1779 For example, the extensive transfer of teleost fishes in aquaculture (Gozlan 1780 et al., 2010; FAO, 2014) has caused widespread decline and even extinctions 1781 of native ichthyofauna, notably in Africa (e.g., Barel et al., 1985; Witte et al., 1782 1992; Canonico et al., 2005). Negative impacts of invasive aliens on a natural 1783 system are often the result of a complex interplay of factors, which may include 1784 direct predation (e.g. Barel et al., 1985; Witte et al., 1992), hybridisation (e.g. 1785 D'Amato et al., 2007; Angienda et al., 2011), and competition for food (e.g. 1786 Ogutu-Ohwayo, 1990) or breeding grounds (Canonico *et al.*, 2005). 1787

The cichlid species *Oreochromis hunteri* (Günther, 1889) is endemic to 1788 Lake Chala (locally 'Challa' after a nearby village; Kenya/Tanzania, Fig. 1.3) 1789 and is currently listed on the IUCN red list as 'critically endangered', primarily 1790 because of its exclusive occurrence in this isolated crater lake (Bayona et al., 1791 2006). Fish fossils in the lake's sediments indicate that this species has been 1792 present for at least 25,000 years (Fig. 1.5), but it may already have colonised 1793 the lake much earlier, not long after the crater basin was formed >250,0001794 years ago (Chapter 3). Despite being the type species of the genus *Oreochromis* 1795 (Günther, 1889), the general ecology of *O. hunteri* has never been studied in 1796 detail. Very little information is available on O. hunteri's breeding and feeding 1797 behaviour (Trewavas, 1983): although Lowe (1955) observed juveniles feeding 1798 on algae and detritus at the shore, nothing is known about the adults, and an 1799 in-depth study of its feeding habits is currently lacking. 1800

Three other cichlid species were introduced to Lake Chala sometime in 1801 the second half of the 20^{th} century, possibly as late as the early 1980s (Dadzie 1802 et al., 1988; Seegers et al., 2003). These include a haplochromine from the 1803 Astatotilapia bloyeti (Sauvage, 1883) complex and two tilapiines, the Redbreast 1804 tilapia (Coptodon rendalli (Boulenger, 1897)) and an unknown Oreochromis 1805 species that we here refer to as *Oreochromis* cf. korogwe, after Chapter 3. These 1806 alien species are surviving well, and especially O. cf. koroque and C. rendalli 1807 may interfere with the endemic O. hunteri, because they are phylogenetically 1808 closely related and have similar ecological preferences. Moreover, the deep open 1809 waters and steep rocky crater walls of Lake Chala provide limited opportunity 1810 for niche segregation. Tilapiine cichlids have in many places outcompeted 1811 native species (Canonico et al., 2005) or hybridised with indigenous tilapia 1812 species (D'Amato et al., 2007; Angienda et al., 2011) due to their tolerance to 1813 a wide range of environmental conditions. Although a combined morphometric 1814 and genetic analysis did not find traces of hybridisation between the introduced 1815 and endemic *Oreochromis* in Lake Chala (Chapter 3), the direct competition 1816 for resources has not yet been assessed. 1817

The trophic versatility of tilapiine cichlids is not only advantageous when coping with new competitors, it is also highly useful in (seasonally) changing environments. Particularly in tropical lakes, wind-induced seasonal upwelling can be an important driving factor of temporary increases in food availability (Lowe-McConnell, 1987), and tilapia species have been described to adapt their feeding behaviour to such events (Spataru, 1976; Man & Hodgkiss, 1977; Spataru, 1978a). In the case of Lake Chala, the short-lived phytoplankton bloom which often develops during the windy mixing season (Barker *et al.*, 2011; Wolff *et al.*, 2011; Buckles *et al.*, 2014) may therefore create seasonal changes in the diet composition of the Lake Chala cichlids, and potentially to a different extent in the indigenous and the introduced species.

In this context, this Chapter aims to document the feeding habits of *Oreochromis hunteri* and the two introduced tilapiine cichlids in Lake Chala, with special emphasis on the link with seasonal fluctuations in resource availability. In this way, we aim to assess the current and potential future impact of the introduced species on the endemic *O. hunteri* population.

1834 4.3 Material and methods

1835 4.3.1 Study system

4

Lake Chala (03°19'S, 37°42'E; Fig. 1.3) is a deep and permanently stratified 1836 crater lake, located at an elevation of \sim 880 m on the border between Kenya 1837 and Tanzania immediately east of Mount Kilimanjaro. Contained within a 1838 steep-sided caldera basin, its rocky shoreline drops down near-vertically to a 1839 depth of ca. 55 m (Opitz, 2008), from where soft fine-grained sediments slope 1840 more gently towards a maximum depth of around 92 m in the middle of the 1841 lake (Moernaut et al., 2010; Fig. 4.1). For most of the year oxygenated water 1842 reaches down to between 15 and 30 m only (Wolff et al., 2011; Wolff et al., 2014), 1843 effectively restricting viable fish habitat to the upper half of the rocky crater 1844 walls and the open-water (limnetic) zone. Submerged rocks are covered with an 1845 algal biofilm and are interspersed with small sandy patches filling up crevices 1846 between them. These patches probably increase in areal extent as they fill up 1847 the rock crevices at water depths beyond wave-induced turbulence. Trapping 1848 the sinking litter of near-shore trees and shrubs, they are likely inhabited by a 1849 diverse zoobenthos community down to where dissolved-oxygen levels become 1850 too low. For most of the year, feeding opportunity for the Lake Chala fishes 1851 is limited to this near-shore habitat, as very low productivity of the limnetic 1852 zone (secchi-disk transparency ranging between 6 and 9 m; van Bree et al., 1853 unpublished data) implies poor feeding conditions. However, combination of 1854 the seasonal minimum in solar insolation with strong monsoon winds causes 1855

intermittent deep water-column mixing down to between 35 and 45 m depth
between late June and September. These short-lived mixing events induce
upwelling of the dissolved nutrients that had been accumulating in deeper
water layers, often (but not always) leading to an extensive phytoplankton
bloom towards the end of the dry season (Barker *et al.*, 2011; Wolff *et al.*,
2011; Buckles *et al.*, 2014). This bloom dramatically enhances limnetic feeding
opportunity for the local fishes.

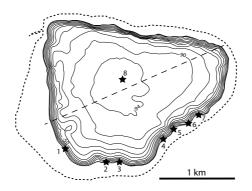


Figure 4.1: Bathymetric map of Lake Chala with indication of food source sampling locations. Location 1-7 represent samples collected from rocks and sandy patches, location 8 is the position where limnetic samples were collected. Dashed line represents the country border between Kenya and Tanzania

1863 4.3.2 Fish and food source collection

Fish samples were collected at monthly intervals over a 20-month period 1864 between February 2014 and October 2015, via purchase from local fishermen 1865 using gill nets suspended from floats and set along the shoreline. Fishermen 1866 adapt the location of their nets to the presence of fish, which fluctuates from 1867 the surface (July – August) to >6m (May – June; November – December). 1868 Fishing in Lake Chala has strongly dwindled over the last decade mainly due 1869 to interference from illegal alcohol distillers, and most fishermen go home to 1870 tend their farms during the seasons when the large tilapiines dwell at depths 1871 beyond the reach of their nets (May – June and November – December). We 1872 nevertheless succeeded to collect fish specimens during most months, yielding 1873

a total of 41 specimens of which 36 contained sufficient gut content for analysis 1874 (Table 4.1). We used a combination of gut content and stable isotope analyses 1875 to document both short and long-term feeding habits of Lake Chala fish. The 1876 guts were immediately dissected to halt further digestion, divided into their 1877 rostral, middle and caudal part, and stored in absolute ethanol (99.8%). For 1878 stable-isotope analysis, dorsal muscle tissue was collected from 17 specimens 1879 using a sterile scalpel, and stored frozen up to the moment of sample processing. 1880 Potential food source samples were obtained in September 2015 by scraping off 1881 the epilithic algal mats from shoreline rocks and by rinsing sand from submerged 1882 sandy patches at 14 different localities on the Kenvan side of the lake (Fig. 4.1). 1883 These samples were frozen for later analyses. Limnetic phytoplankton samples 1884 were collected by filling 100 mL bottles with surface water, monthly between 1885 February 2014 and October 2015, and at two additional depth intervals (5 1886 and 10 m) between January 2014 and January 2015 using a UWITEC water 1887 sampler. Immediately after collection these water samples were fixed in situ 1888 with an alkaline Lugol's solution and formalin. 1889

1890 4.3.3 Laboratory methods

The abundance of gut content was determined qualitatively upon return to 1891 Ghent University, Belgium. Guts that were filled over the entire length of 1892 the rostral part were considered full, whereas guts with large empty gaps but 1893 still sufficient material for counting were considered half-full. All others were 1894 considered empty for the purpose of this study. The middle and caudal parts of 1895 the guts contained strongly digested material only and were not further analysed. 1896 The content of the first cm of each gut was transferred to a 2-mL Eppendorf 1897 vial and diluted with distilled water. If the first cm did not contain material, 1898 which sometimes occurred in specimens with half-full guts, the first food bolus 1899 encountered within the anterior 10 cm of the guts was used. Littoral food-1900 source samples were similarly diluted in distilled water. One mL of the resulting 1901 suspension was pipetted into a sedimentation chamber of 10 mL and left to 1902 settle down. The limnetic phytoplankton samples were analysed following the 1903 Utermöhl method (Utermöhl, 1931, 1958) and contained Chlorophyta (green 1904 algae), Bacillariophyta (diatoms), Dinophyta, Euglenophyta, Chrysophyta. 1905 Cryptophyta and Cyanobacteria (blue-green algae) in variable quantities and 1906 proportions. In each gut and littoral sample 200 algal/cyanobacterial cells, 1907

Table 4.1: Overview of fish specimens available for gut analysis, collected in Lake Chala between February 2014 and September 2015. Values indicate the number of collected specimens per month and per species, whereas cell colour represents the abundance of gut content: dark grey cells indicate full guts, light grey cells indicate half-full guts and white cells indicate empty guts

Month	O. hunteri	O. cf. korogwe	C. rendalli	Total
February 2014	2	1	1	4
March 2014	1	1	1	3
April 2014	1	1	1	3
May 2014	1	1	1	3
June 2014	-	-	-	-
July 2014	1	1	-	2
August 2014	1	1	-	2
September 2014	1	1	1	3
October 2014	-	-	-	-
November 2014	-	-	-	-
December 2014	1	1	1	3
January 2015	1	-	1	2
February 2015	1	-	1	2
March 2015	-	-	-	-
April 2015	1	-	1	2
May 2015	-	-	-	-
June 2015	1	-	1	2
July 2015	-	-	-	-
August 2015	1	-	-	1
September 2015	4	1	1	9
Total	18	9	14	41

colonies or detritus items were counted; in the phytoplankton samples at 1908 least 500 solitary cells or colonies. Enumeration and taxonomic identification 1909 was done using a Kyowa inverted microscope for gut and littoral samples 1910 (Ghent University), and an Olympus CKX 41 inverted microscope for the 1911 phytoplankton (Meise Botanic Garden). All analyses were done with a 100x 1912 immersion oil objective. Different algal/cyanobacterial taxa made up the 1913 bulk of most studied samples, and were identified down to genus level for the 1914 dominant taxa and down to order level for other taxa (31 in total, cf. below). 1915 Colonies were counted as one item, but the number of individual cells in each 1916 colony was also recorded. The average number of cells per colony was compared 1917 between guts and food sources for each colonial taxon. If this average number 1918

of cells differed, for example due to rapid disintegration of colonial genera in the fish gut, colony counts were rescaled accordingly. Plant remains were classified as detritus; most of these must originate from terrestrial plant litter because submerged or emergent aquatic macrophytes are almost absent in Lake Chala. The loose sponge spicules that were occasionally found were also regarded as detritus, rather than evidence for direct consumption of sponges.

For stable-isotopic analysis of fish tissue, samples were dried overnight at 60°C and pulverised using a mortar. Subsamples were weighed to the nearest 0.01 mg before being put into tin capsules (8x5 mm). Replicates of the same sample were analysed for specimens of *O. hunteri* and *C. rendalli*. Stable isotope analyses were carried out at the Isotype Bioscience Laboratory (ISOFYS) of Ghent University. The samples were run for nitrogen (N) and carbon (C) isotope ratios, which are expressed in delta notation:

¹⁹³²
$$\delta^{15}$$
N or δ^{13} C = [(Rsample / Rstandard) 1] x 1000

with $R = {}^{15}N/{}^{14}N$ or ${}^{13}C/{}^{12}C$ relative to a standard, which is atmospheric nitrogen for N and Pee Dee Belemnite limestone for C.

1935 4.3.4 Data analyses

Further data analyses were performed in R version 3.3.2 (R Development Core 1936 Team, 2016). Species accumulation curves (package 'vegan' v.2.4-3; Oksanen 1937 et al., 2017) were used to establish that 200 algal/cyanobacterial cells, colonies 1938 or detritus items were sufficient to capture the composition of ingested food. 1939 and that enough fish specimens were analysed for comprehensive assessment of 1940 fish diets in Lake Chala. To assess seasonal differences in gut content, fishes of 1941 each of the three cichlid species considered were initially analysed collectively 1942 per season. Guts of fishes caught during the mixing season (July to September) 1943 were assembled into one group, and those collected during the predominantly 1944 stratified season (October to June) into another group, per consecutive year. 1945 Our gut collections during 2014 and 2015 hence resulted in observations from 1946 four seasons, for which diet composition was compared with multivariate 1947 techniques using the packages 'vegan' and 'MASS' v.7.3.45 (Venables & Ripley, 1948 2002). A Bray-Curtis similarity matrix was generated based on square-root 1949 transformed counts of the gut contents. We compared the multivariate spread 1950 between samples, as homogeneity of spread is an important assumption for 1951

permutational multivariate analysis of variance (PERMANOVA). We tested for
differences in seasonal diet using PERMANOVA (999 permutations), followed
by pairwise comparison *post-hoc* tests. Results were visualised using non-metric
multidimensional scaling (NMDS).

To compare observed seasonal differences in diet composition with seasonal 1956 variation in food sources, we related the gut content of the three fish species 1957 combined per consecutive month to the limnetic and littoral food-source samples. 1958 Counts were summed per month and divided by the total number of food items 1959 in that month to calculate a percentage by number (%N; after Hyslop, 1980). 1960 A similar method was applied to estimate the relative importance of different 1961 algal/cyanobacterial taxa in the phytoplankton samples, by combining samples 1962 from 0, 5 and 10 m depth when applicable, into the percentage by number 1963 (%N) per month. Littoral samples were only available for September 2015, 1964 hence %N for rock scrapes and sandy patches was calculated from the overall 1965 average of the 14 sampled localities. All the above-mentioned metrics were 1966 calculated both at the phylum level (including the categories of detritus and 1967 sand grains), and at the highest taxonomic resolution feasible. The results 1968 were visualised and compared using bar plots. 1969

Generalised (i.e., long-term average) interspecific diet differences were 1970 assessed using δ^{15} N and δ^{13} C values. A possible influence of high lipid values 1971 on δ^{13} C results can be controlled for by simultaneous lipid extraction or 1972 mathematical normalisation. Animal lipid content has a strong relationship 1973 with the C:N ratio, and a C:N ratio <3.5 (corresponding to a 5% lipid content) 1974 is preferred (Post *et al.*, 2007). None of our fish specimens crossed that 1975 threshold, hence normalisation of the δ^{13} C values was deemed unnecessary. 1976 ANOVA followed by *post-hoc* Tukey-tests was used to test the significance of 1977 differences in δ^{15} N and δ^{13} C values between the fish species. We also assessed 1978 diet disparity (i.e. isotopic ranges) for each species by calculating the mean 1979 Euclidean distance of each point in the δ^{13} C: δ^{15} N plot to its group centroid, 1980 and using the sum of variances on the two axes. 1981

Gut composition between species was compared using PERMANOVA after the assumption of homogeneity was checked. We used the same similarity matrix as for seasonal comparisons and visualised the results using NMDS. We then determined the dietary importance of food items per fish species and per season using percentage by number (%N) and visualised the results ¹⁹⁸⁷ in bar plots to assess the exact diet differences between the respective fish ¹⁹⁸⁸ species. Although ontogenetic diet shifts are widely reported in tilapiine species ¹⁹⁸⁹ (Ribbink, 1990; Piet, 1998; Njiru *et al.*, 2004), we found no influence of body ¹⁹⁹⁰ size on the diet of the studied species (Supplementary Information C).

1991 4.4 Results

4

¹⁹⁹² 4.4.1 Seasonal differences in gut contents and food sources

The guts of 18 Oreochromis hunteri, 9 O. cf. koroque and 14 Coptodon rendalli 1993 from Lake Chala together contained 28 different prey types identified at the 1994 lowest-feasible taxonomic level (Supplementary Table C.1), including two types 1995 of cyanobacteria, 24 types of algae, detritus and sand grains. Unicellular algae 1996 and cyanobacteria make up between 45 and 95% of the analysed gut contents. 1997 The 29 samples of limnetic phytoplankton of Lake Chala contained a total of 27 1998 different prey types, namely three types of cyanobacteria and 24 types of algae 1999 (Supplementary Table C.2). The 12 littoral rock samples and 2 littoral sand 2000 samples contained 24 and 18 different prey types, respectively (Supplementary 2001 Table C.3). . 2002

Table 4.2: Bonferroni-corrected p-values of i) *post-hoc* homogeneity Tukey-tests per season (above diagonal) and ii) of *post-hoc* PERMANOVA pairwise comparisons of Bray-Curtis similarity matrices per season (below diagonal). The asterisks (*) signify significance at the adjusted 5% level.

	Feb '14 - Jun '14	Jul '14 - Sep '14	Dec '14 - Jun '15	Sep '15
Feb '14 - Jun '14		0.91	0.16	0.24
Jul '14 - Sep '14	0.01^{*}		0.54	0.07
Dec '14 - Jun '15	0.47	$< 0.01^{*}$		$< 0.01^{*}$
Sep '15	0.09	0.02^{*}	$< 0.01^{*}$	

A two-dimensional NMDS plot was generated with acceptable stress values (0.16) and hence represents the data well. This plot (Fig. 4.2) shows that gut content composition varies greatly between seasons, whereas interspecific differences are less pronounced. We found sample homogeneity to differ significantly between seasons (F = 6.36, p < 0.01), but *post-hoc* Tukey tests revealed only one significant pairwise distinction (between the two seasons of 2015; Table 4.2). Therefore, assumptions are met for the majority of inter-season comparisons and PERMANOVA can be safely executed, although the results of one comparison should be treated with caution. Overall, we found a highly significant impact of seasonality on diet composition (F = 4.63, p < 0.01), and *post-hoc* testing revealed that all pairwise differences between mixing and predominantly stratified periods were borderline to highly significant on the Bonferroni-corrected 5% confidence interval (Table 4.2).

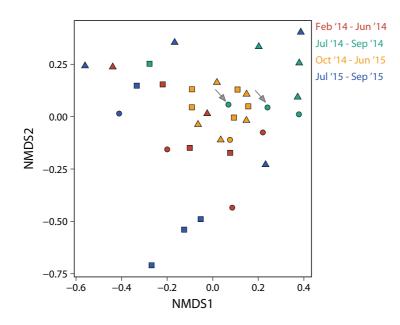


Figure 4.2: Non-metric multidimensional scaling ordination (NMDS; stress = 0.16) visualising similarity in gut content. Shapes represent interspecific differences between the endemic *O. hunteri* (\triangle) and introduced *O.* cf. *korogwe* (\circ) and *C. rendalli* (\square). Colour describes seasonality, in which a distinction is made between predominantly stratified seasons (2014, red; 2015, yellow) and mixing seasons (2014, green; 2015, blue). Arrows indicate specimens of *O.* cf. *korogwe* collected in July and August 2014.

To comprehend the sources of seasonal variation in gut contents, we 2016 compared the relative abundances of different food types in the fishes' guts 2017 with those of the limnetic and littoral food sources, both at the phylum level 2018 (Fig. 4.3) and with a higher taxonomic resolution (Fig. 4.4). Apart from 2019 February 2014, the proportion of detritus and sand grains in the guts is much 2020 lower during the during the mixing season (6-13%) than during the rest of the 2021 year (27-53%; Fig. 4.3b). Whereas all phyla of algae/cyanobacteria, except 2022 Chrysophyta, were found to occur in both the limit and littoral zone, the 2023

detritus and sand in the fish guts have a littoral origin, as these non-algal food 2024 types are (predictably) absent in the open surface water but represent almost 2025 half of the littoral food sources (Fig. 4.3c). Chlorophyta are abundant in the 2026 gut contents at the start of the 2014 mixing season (from 10% to 30-50%). 2027 which coincides with high abundances of the chlorophyte *Tetraedron* Kützing 2028 (1845) in the open water column (Fig. 4.4). Indeed, *Tetraedron* is during that 2029 period abundant in the fish guts. It also occurs in large proportions in the guts 2030 in certain months of the predominantly stratified season, but is in those months 2031 less abundant in the limnetic zone. The proportion of Bacillariophyta (diatoms) 2032 in the guts remains comparatively stable throughout the year, but composition 2033 of diatom genera is highly variable. Nitzschia Hassall (1845) and Staurosirella 2034 Williams & Round (1987) represent a significant part of the gut contents 2035 during the predominantly stratified season (10-45% and 5-18%, respectively). 2036 Nitzschia makes up a considerable part of littoral food sources (15-30%), and 2037 although Staurosirella represents only a small fraction of the littoral samples 2038 $(\sim 0.5\%)$, it requires a substrate for growing, arguing for a littoral habitat. 2039 Staurosirella also occurs in the limnetic zone during the mixing season, when 2040 zooplankton abundance peaks in response to the phytoplankton bloom and 2041 their faecal pellets provide abundant substrates for this diatom. Simultaneously, 2042 however, *Staurosirella* abundance in fish guts strongly decreases (0-1%), and it 2043 is replaced by *Fragilaria* s.l. Lyngby (1819) in July and August (12-28%) and 2044 Afrocymbella Krammer (2003), more specifically the endemic species A. barkeri 2045 Cocquyt & Ryken (2016), in September (11%). These genera also successively 2046 peak in the limnetic phytoplankton at that time. Notably, the Chlorophyte 2047 Treubaria Bernard (1908) is highly abundant in the water column in July 2014, 2048 but is rare in the guts of fish collected in that month. 2049

The expected similar trend during the mixing season of 2015 was not re-2050 alised. Instead, gut contents from September 2015 largely consist of Cyanobac-2051 teria, although the gut of one specimen of O. hunteri was filled with green algae. 2052 All guts from specimens collected at that time have at least some content, 2053 but in most guts food boluses are interspersed with large stretches of empty 2054 gut (Table 4.1). Absolute abundances of the major algal/cyanobacterial taxa 2055 present in the water column (Fig. 4.5) show that a seasonal diatom and green 2056 algae bloom was mostly absent that year. Other important seasonal abundance 2057 shifts that take place in the limit phytoplankton are not reflected in the fish 2058

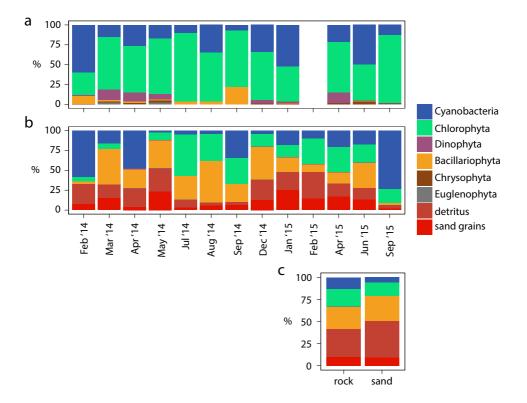


Figure 4.3: Proportion (%N) per month of algal/bacterial phyla in limnetic phytoplankton (a), and algal/bacterial phyla, detritus and sand in fish guts (b) and littoral food sources (c). 'Rock' refers to samples scraped from rocks; 'sand' refers to patches of sand.

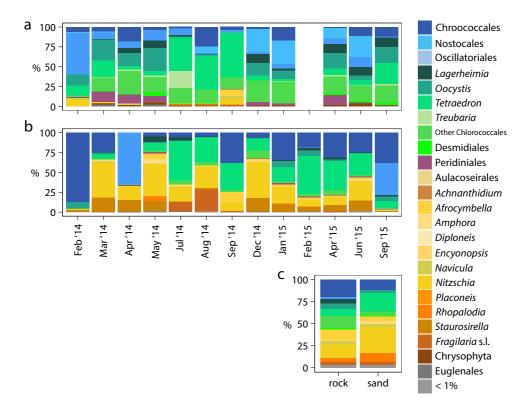


Figure 4.4: Proportion (%N) per month of algal/bacterial genera and orders in the limnetic phytoplankton (a), fish guts (b) and littoral food sources (c).

genus *Oocystis* Braun (1909) from March to May, and increased numbers of Nostocales Borzi (1914) (Cyanobacteria), mainly *Cylindrospermopsis* Seenayya & Subbarayu (1972), in December and January. Overall, gut samples taken during the mixing season resemble the limnetic phytoplankton composition at that time, whereas gut samples taken at different times of the year contain higher littoral contents.

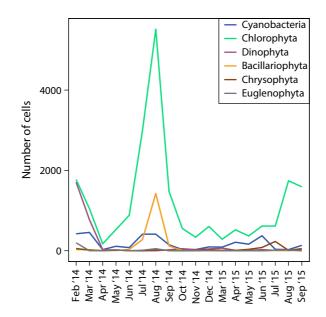


Figure 4.5: Total abundance (counts in 100 mL surface water) of cells or colonies of the six most common algal/bacterial phyla in Lake Chala, during 20 consecutive months from February 2014 to September 2015.

²⁰⁶⁶ 4.4.2 Interspecific and size-related differences

There was no significant difference in the homogeneity of gut contents grouped per fish species (F = 0.31, p = 0.73), hence assumptions for PERMANOVA are fulfilled. Although PERMANOVA indicated a borderline significant difference in gut content composition between fish species (F = 1.90, p = 0.05), *post-hoc* tests revealed that none of the pairwise differences were significant at the Bonferroni-adjusted 5% significance level (Table 4.3). The available sample sizes did not allow us to simultaneously test for the influence of both species

and season, but bar plots taking into account both factors (Fig. 4.6) show some 2074 interesting interspecific differences. Although all species show the seasonal 2075 shift between littoral and limnetic food sources (Fig. 4.6a), during the mixing 2076 season of 2014 the diet of O. hunteri has a much higher contribution of the 2077 blooming limnetic algae *Tetraedron* and *Fragilaria* s.l. (Fig. 4.6b). In contrast, 2078 O. cf. korogwe specimens caught during that time (particularly those from 2079 July and August 2014) show diet compositions similar to those observed in 2080 all species during the predominantly stratified season. This is also reflected 2081 in the NMDS plot (Fig. 4.2), where O. cf. korogwe specimens from July and 2082 August cluster with specimens collected in the stratified season. Only the 2083 individual caught in September 2014 contains high abundances of the limnetic 2084 algae blooming at that time, such as Afrocymbella and Tetraedron. The gut 2085 content of the single C. rendalli specimen obtained during the mixing season of 2086 2014 consisted mostly of Cyanobacteria. Some minor interspecific differences 2087 can also be observed during the predominantly stratified season, for example 2088 that O. cf. koroque has higher abundances of the diatom Staurosirella in its 2089 gut than the other two cichlids. 2090

Table 4.3: Bonferroni-adjusted p-levels of pairwise comparison on Bray-Curtis similarity matrices of gut composition per species (*post-hoc* PERMANOVA).

	O. hunteri	O. cf. korogwe
O. cf. korogwe	0.22	
C. rendalli	0.18	1.00

The total isotopic range in all Lake Chala tilapiines combined is higher 2091 for δ^{13} C (6.18) than for δ^{15} N (2.88). Muscle tissue of the three fish species 2092 differed significantly in stable-isotopic composition (Fig. 4.7), both for $\delta^{15}N$ (F 2093 = 19.12, p < 0.01) and somewhat less for δ^{13} C (F = 4.5, p = 0.03). Average 2094 δ^{15} N was higher for O. hunteri ($\overline{x} = 10.11$) and C. rendalli ($\overline{x} = 10.22$) than 2095 for O. cf. koroque ($\overline{x} = 8.02$), whereas O. hunteri ($\overline{x} = -24.60$) was on average 2096 more depleted in δ^{13} C than both O. cf. korogwe ($\overline{x} = -21.84$) and C. rendalli 2097 $(\overline{x} = -22.76)$. Disparity metrics indicate that isotopic disparity is higher in O. 2098 hunteri than in the two introduced species (Table 4.4), particularly in δ^{13} C 2099 where its range approaches the total observed range (Fig. 4.7). 2100

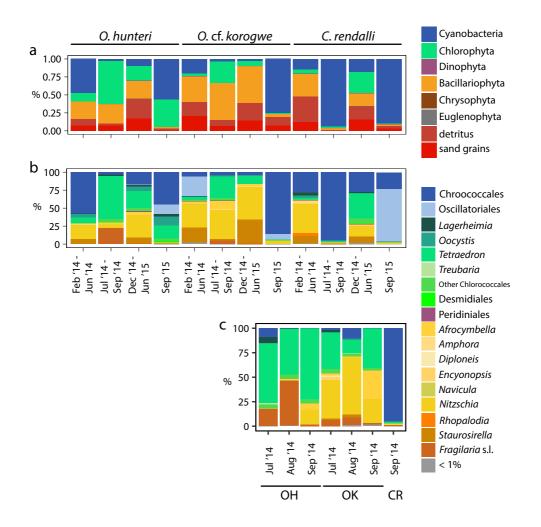


Figure 4.6: Proportion (%N) of algal phyla, detritus and sand (a); and algal genera and orders (b) in fish gut content per season and per fish species. Panel c displays the gut content per month and per species during the annual period of phytoplankton bloom in 2014

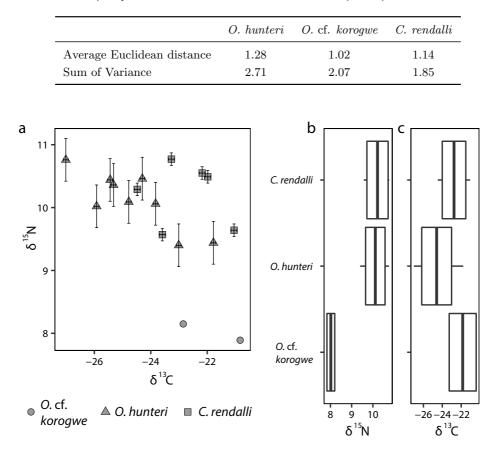


Table 4.4: Disparity measures of C:N-ratios in the three tilapiine species of Lake Chala.

Figure 4.7: a Carbon and nitrogen stable-isotope bi-plot with symbols representing the endemic *O. hunteri* (yellow \triangle) and introduced *O.* cf. *korogwe* (green \circ) and *C. rendalli* (red \Box). Error bars represent the standard deviation in stable-isotopic composition of two replicate samples of the same specimen, which were only available for *O. hunteri* and *C. rendalli*. Box plots display mean, standard deviation (SD) and minimum and maximum values per species, for δ^{15} N (**b**) and δ^{13} C (**c**).

2101 4.5 Discussion

Apart from some periods during the predominantly stratified season, when fish reside somewhat deeper and stay very close to shore, and are hence more difficult to catch, specimens were collected in 14 out of 20 months (Table 4.1). Sample sizes are relatively modest; nevertheless, our study is the first to provide insights into the feeding habits of both the endemic *O. hunteri* and introduced *O.* cf. *korogwe* and *C. rendalli* in Lake Chala, although we wish to

be prudent with certain interpretations. In the examined guts we did not find 2108 evidence of live consumption of animals, not from the rocks nor benthic fauna 2109 nor pelagic zooplankton. Similarly, the isotopes do not give indication of the 2110 high mean trophic level occupied by omnivores. The results of our gut analyses 2111 therefore indicate that all three species currently inhabiting Lake Chala retain 2112 a herbivorous diet supplemented with detritus throughout the year, like most 2113 other tilapias (Trewavas, 1983). However, we found highly significant seasonal 2114 shifts in the feeding behaviour of all three species, that were much larger 2115 than the observed interspecific differences. These shifts are strongly linked 2116 to seasonal fluctuations in phytoplankton abundance, highlighting the role of 2117 resource availability in foraging behaviour. 2118

The three cichlids appear to be restricted to exclusively littoral food sources 2119 when the water column is predominantly stratified (February – June 2014 and 2120 October – June 2015), but the extensive bloom of green algae (Chlorophyta) and 2121 diatoms (Bacillariophyta) during the mixing season (July – September) provides 2122 plenty of food to temporarily draw fish towards the open water, where they 2123 feed on the successive peaks of different algal taxa (Fig. 4.4). Remarkably, the 2124 short-lived yet considerable limnetic Dinophyta bloom, generally taking place 2125 from February to March (at least in 2014; Fig. 4.5), is not exploited, indicating 2126 that Lake Chala cichlids prefer to retain their littoral feeding habits over 2127 abundant ingestion of Dinophyta. Also, the Chlorophyte Treubaria is largely 2128 absent from the guts in July 2014, although it is at that time highly abundant 2129 in the water column (Fig. 4.4). Other planktivorous Oreochromis species use 2130 their pharyngeal jaws to mechanically process phytoplankton (Greenwood, 2131 1965; Trewavas, 1983), and potentially, the rather large Treubaria (~ 40 μ m, 2132 including the spines) is broken beyond recognition by these jaws before it enters 2133 the gut. In contrast, the strong siliceous values of diatoms may enable them to 2134 pass these jaws relatively unscathed. 2135

We detected crucial differences in gut content composition between mixing seasons of subsequent years. The lack of an extensive phytoplankton bloom in 2015 (Fig. 4.5) is reflected in the guts by high proportions of Cyanobacteria (Fig. 4.4) yet low food abundance (Table 4.1), indicating a decrease in food consumption. Although some littoral material (detritus, sand grains) is present in the guts from these months, it is not abundant, suggesting that fish venture out into open waters despite sub-optimal feeding conditions. The presence of some individuals with guts filled with green algae demonstrates that abundant
phytoplankton can be locally available, and may still trigger fish to migrate to
the limnetic zone.

4

Importantly, similar seasonal shifts in feeding habits were observed for 2146 all species (Fig. 4.6a), indicating that the introduced species have rapidly 2147 learned to exploit the seasonal productivity patterns characteristic for Lake 2148 Chala. Although this observation is not sufficient to directly demonstrate 2149 competition between species (Zengeya et al., 2011), it provides strong circum-2150 stantial evidence that O. hunteri is presented with serious competition for food. 2151 Indeed, we could not find significant interspecific differences in year-round gut 2152 content. However, substantial distinctions between species are present in the 2153 isotopic signal (Fig. 4.7), with lower δ^{13} C, higher δ^{15} N and higher isotopic 2154 disparity in O. hunteri than in its non-native counterparts, suggesting at least 2155 some form of resource partitioning (Zengeva et al., 2011). The distinction is 2156 mostly reflected in the guts of specimens collected during the mixing season, 2157 potentially indicating interspecific segregation during the phytoplankton bloom 2158 only. Indeed, whereas O. hunteri guts contain limnetic phytoplankton (Fraq-2159 *ilaria* s.l., *Tetraedron*) from July onwards, O. cf. *koroqwe* appears to linger 2160 in the littoral until Afrocymbella peaks in September (Fig. 4.6b). The one 2161 specimen of C. rendalli collected during this period had only Cyanobacteria in 2162 its gut, hence this species potentially does not feed on the green algae/diatom 2163 bloom at all. Phytoplankton is on average more depleted in δ^{13} C than epilithic 2164 algae and detritus (Finlay & Kendall, 2007). As regards Lake Chala, Hurrell 2165 et al. (2011) found very low δ^{13} C values for the blooming diatom genera 2166 Nitzschia and Afrocymbella. The longer exposure of the endemic O. hunteri to 2167 phytoplankton could explain the more depleted δ^{13} C values, and its enhanced 2168 isotopic disparity. 2169

Gut content composition and stable isotopes thus indicate that the in-2170 digenous O. hunteri is taking more advantage of the opportunities provided 2171 by the annual phytoplankton bloom in Lake Chala than introduced cichlids. 2172 The abundance of easily accessible food present in the limit zone at that 2173 time is likely preferable over scarcer littoral food sources that require active 2174 exploration and competition with other individuals. The introduced tilapiines 2175 seemingly have not (yet) synchronised their behaviour to the seasonal pat-2176 terns in food availability in Lake Chala, whereas the presence of O. hunteri 2177

dates back at least 25,000 years (Chapter 6), which has resulted in a perfect 2178 synchrony between the endemic species and lake dynamics. Therefore, O. 2179 hunteri potentially has a competitive advantage over the recent newcomers. 2180 Nevertheless, this advantage strongly depends on the annual recurrence of a 2181 strong bloom of diatoms and green algae in the limnetic zone. The strong 2182 effect of low phytoplankton abundances on gut content during the mixing 2183 season of 2015 therefore raises concerns about the effects of predicted climate 2184 warming in East Africa. Wolff et al. (2011) highlighted the consequences 2185 of climate variability on the strength of the annual phytoplankton bloom in 2186 Lake Chala, as wetter, less windy years can strongly reduce its intensity. The 2187 prognoses for a warmer climate in East Africa include higher rainfall and 2188 higher rainfall variability (Held *et al.*, 2006) and could thus severely impact the 2189 intensity of the phytoplankton bloom. Although O. hunteri has coped with 2190 severe climate-driven habitat changes in its past (Moernaut et al., 2010), in 2191 all likelihood the species has never had to deal with competition from other 2192 species during such events (Chapter6). Climate change could therefore affect 2193 the faith of endemic and introduced fish species in Lake Chala. 2194

Chapter 5

2196	
2197	Distribution and burial of fish remains in near-shore sediments
2198	of Lake Chala
2199	

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2207 5.1 Abstract

The isolated crater lake Chala (Kenya/Tanzania) has a continuous offshore 2208 sediment record containing abundant fossil fish remains, yet its only endemic 2209 cichlid, Oreochromis hunteri, spends most of its life close to the shore. This 2210 raises questions about how fish fossils are transported to and eventually buried 2211 in the depositional centre of the lake. Here, we analyse the distribution of fish 2212 fossils (teeth, bones, scales) as well as the remains of diverse other lacustrine 2213 biota throughout a short core spanning the last ~ 180 years, collected near the 2214 shore of Lake Chala, in order to gain insights into sedimentation processes 2215 influencing their burial and preservation. This core displays a strong top-2216 to-bottom gradient of sediment characteristics, from soft, fine-grained and 2217 finely laminated sediments to consolidated basal muds containing sand and 2218 gravel-sized rock particles. We found that the median depth of occurrence 2219 of each type of fossil is strongly related to its hydrostatic properties: fossils 2220 with low specific density are more abundant in the fine-grained upper core 2221 section, whereas fossils with high specific density and/or compact shape are 2222 concentrated in the coarse-grained lower section. This suggests that the lower 2223 core section is equivalent to a lag deposit, predominantly containing those 2224 types of fossils that are least susceptible to the winnowing effect of repeated 2225 re-suspension and transport towards greater water depths offshore. In the 2226 calm bottom waters of meromictic Lake Chala, not many fish bones are likely 2227 to be transported over long distances before final burial. We therefore argue 2228 that fish teeth preserved in the offshore sediment record have predominantly 2229 been deposited in situ. Although we observed a significant correlation between 2230 fossil teeth abundance and median grain size, sediment properties did not 2231 have a strong influence on tooth size or fractional abundance of teeth from 2232 different jaws. Considering further that the offshore sediments of Lake Chala 2233 are uniformly fine-grained and laminated, at least throughout the last 25,000 2234 years, we suggest that any changes in sedimentation dynamics that may have 2235 occurred during that time must have had a negligible influence on the fish 2236 fossil assemblages buried at Lake Chala's depositional centre. 2237

2238 5.2 Introduction

For several decades now, cichlid fishes have been under study as a model system 2239 for evolution. Rapid adaptation of the trophic apparatus to a wide variety 2240 of available niches, coupled to sexual selection and subsequent reproductive 2241 isolation, has been considered pivotal in the origin of hundreds of species in 2242 what is a prime example of adaptive radiation (e.g., Stiassny & Meyer, 1999; 2243 Danley & Kocher, 2001; Salzburger et al., 2014). Morphological change in 2244 the cichlid's trophic apparatus in response to environmental pressures has 2245 been described to occur on decadal timescales (Witte *et al.*, 2008; van Rijssel 2246 et al., 2015), and the plasticity of this apparatus is often suggested to have 2247 influenced speciation (e.g., Meyer, 1990; Muschick et al., 2011; Gunter et al., 2248 2013). Yet the absence of a good fossil record for this group has hampered the 2249 study of morphological change through time beyond historical records (Murray, 2250 2001). In this context, fossil fish remains preserved in continuous lake-sediment 2251 records could fill an important knowledge gap. Especially fossil fish teeth can 2252 potentially be used to trace trophic adaptation to local habitat changes over 2253 very long timescales. The cichlid fossil archives available in sediment sequences 2254 recovered through lake drilling are thus starting to be explored (Reinthal 2255 et al., 2011). These cores are typically collected in deepwater areas offshore 2256 to guarantee continuity in sediment deposition (Wilke *et al.*, 2016). However, 2257 in many lakes the majority of cichlid species live close to shore (e.g., Fryer, 2258 1959b; Ribbink et al., 1983), raising the question of how their remains reach 2259 their final burial place offshore. 2260

This issue is also relevant to the fish fossil record of Lake Chala (Kenya/-2261 Tanzania). Abundant fossil cichlid teeth have been recovered from a 25,000-year 2262 sediment sequence collected at Lake Chala's depositional centre in the middle 2263 of the lake (Chapter 2 & 6), yet gut contents of the endemic cichlid Oreochromis 2264 hunteri (Günther, 1889) indicate that this species feeds on littoral food sources 2265 for most of the year, and thus must typically reside close to shore (Chapter 4). 2266 Understanding how the fossil remains of this species end up in the middle of 2267 the lake is an important prerequisite for correct interpretation of the recovered 2268 fossil assemblages. 2269

Lake Chala is a 92-meter deep, steep-sided crater lake at the foot of Mount Kilimanjaro with a uniform basin morphometry (Fig. 5.1). Its near-vertical

rocky crater walls drop down to a depth of 50-55 m, from where soft sediments 2272 slope down more gently to approximately 92 m at the depositional centre 2273 (Opitz, 2008; Moernaut et al., 2010). For most of the year, oxygenation of 2274 the water column is limited to a depth of 15-20 m, but cooling of the water 2275 surface from late June to September promotes convective mixing down to 2276 between 35 and 45 m (Buckles et al., 2014; De Crop et al., unpublished data). 2277 However, considering the strong control of inter-annual climate variability 2278 on water-column mixing in Lake Chala (Wolff *et al.*, 2011), extreme climate 2279 anomalies occurring once every several decades (e.g., Buckles et al., 2016) may 2280 induce mixing to even greater depths, reaching down to soft-sediment bottom 2281 areas near the lake periphery. Seismic-reflection data revealed that pronounced 2282 lake-level fluctuations have occurred since Lake Chala's origin approximately 2283 250,000 years ago (Moernaut et al., 2010; Verschuren et al., 2017), but although 2284 these were accompanied by slightly altered patterns of sedimentation, at least 2285 in the past 25,000 years they were never severe enough to expose near-shore soft 2286 sediments. Also, the depositional centre of Lake Chala has likely always been 2287 anoxic and devoid of a zoobenthos community (Meyer et al., 2018), reflected 2288 in the absence of bioturbation implied by the prevalence of seasonal-scale fine 2289 lamination throughout the sediment sequence (Verschuren et al., 2009; Wolff 2290 et al., 2011). Therefore, overall sedimentation dynamics in these profundal 2291 bottom areas must have remained relatively stable through time. 2292

There are three possible ways in which cichlid teeth may end up at the 2293 centre of Lake Chala. Oreochromis hunteri displays seasonal migration into 2294 the limnetic zone in response to an annual phytoplankton bloom (Chapter 4). 2295 Although fish would not have to venture far into open waters to feed off this 2296 bloom, they may occasionally do so, and fossils could during this period be 2297 deposited in situ. Such in situ deposition may, firstly, be the result of cichlids 2298 continuously replacing their teeth throughout their lives (polyphyodonty; e.g., 2299 Streelman et al., 2003), and individual teeth in offshore sediments may thus 2300 result from shedding events at that location. Secondly, they could represent 2301 remains after death in various stages of disintegration (Elder & Smith, 1988). 2302 Thirdly, fossil teeth are largely deposited close to the lake periphery, and may 2303 subsequently be transported from the shore to the centre through repeated 2304 resuspension and sediment-focussing processes. Occasional extreme deep-2305 mixing events (on decadal and longer time scales) may distort the upper layers 2306

of soft sediment and induce focussing towards the depositional centre. Cichlid fossils that are present in these upper layers may thus be transported further and further from shore until they reach their final burial place in the middle of the lake. Importantly, the likelihood of such extreme events increases during lake lowstands, potentially increasing the proportion of teeth (relative to the total amount of teeth produced by the entire lake population in a given period of time) that is eventually deposited in the middle of the lake.

To shed light on some of the processes involved, in this Chapter we study 2314 the distribution of fish fossils (teeth, bones and scales) as well as the remains 2315 of diverse other lacustrine biota throughout a short surface core collected 2316 near-shore. This way, we aim to determine which of the abovementioned 2317 processes is more likely to influence fish fossil distributions in sediments from 2318 the centre of Lake Chala. Whereas the monotonous sediment composition at 2319 the depositional centre hampers correlation of fossils with sediment properties. 2320 the variable nature of these near-shore sediments facilitates comparison between 2321 fossil distributions and sedimentology. We thus intend to gain insight into 2322 potential influences of altered sedimentation patterns on the composition of 2323 fossil assemblages during lake lowstands. 2324

2325 5.3 Material and methods

2326 5.3.1 Core collection

A 27 cm-long gravity core (CH16-34G) was collected near the shore of Lake Chala (Fig. 5.1; 53 m depth) in November 2016 using a large-diameter UWITEC gravity corer (ϕ 9cm). To preserve the original gradient of water content and degree of sediment compaction the core was extruded upright in the field and stored in Whirlpack bags for transport to the lab in Ghent, Belgium. The upper 16 cm was extruded in 2-cm intervals, whereas the lower 11 cm was extruded in 1-cm intervals, resulting in a total of 19 sediment intervals.

Finely laminated sediments occur in approximately the top 9 cm of the core (Fig. 5.2a). A detailed count of the number of years represented by these laminated layers cannot be obtained from our extruded core. However, correlation of photographs from the freshly-collected core (Fig. 5.2a) with splitcore photographs of partly dewatered surface cores also collected in November 2016 (Fig. 5.2b) in the same near-shore area and at comparable water depths

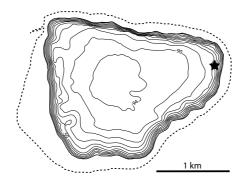


Figure 5.1: Bathymetry of Lake Chala, with indication of near-shore coring location at 53-m water depth

(60 and 65 m; Fig. 5.2b & c), indicates that the 9 cm of laminated sediment were deposited over a time span of approximately 25 years. Extrapolation of the obtained sediment accumulation rate, corrected for the down-core decrease in water content and thus compaction (Fig. 5.3c), indicates that the age of the basal interval of core CH16-34G may be on the order of 180 years (but see below).

2346 5.3.2 Assessment of sediment properties

We used sequential Loss-On-Ignition (LOI) to estimate the organic and car-2347 bonate content for each interval (Heiri *et al.*, 2001). We transferred 1 mL of 2348 wet sediment from each interval into a crucible and assessed weight loss after 2349 14 h at 105°C to determine its H_2O content and porosity (H_2O content by 2350 volume). Organic matter (OM) content was determined in a similar fashion 2351 after burning for 4 h at 550°C, and carbonate content (CaCO₃) was determined 2352 after ashing for 2 h at 1000°C. The rest fraction (calculated by subtracting the 2353 percentages of OM and $CaCO_3$ from 100) is here considered to exclusively rep-2354 resent siliciclastic material, but in reality it also includes an unknown fraction 2355 of biogenic silica derived from diatoms and sponge spicules. 2356

Determining the grain-size distribution of the siliciclastic sediment component requires removal of all carbonate and biogenic material from the sediment,

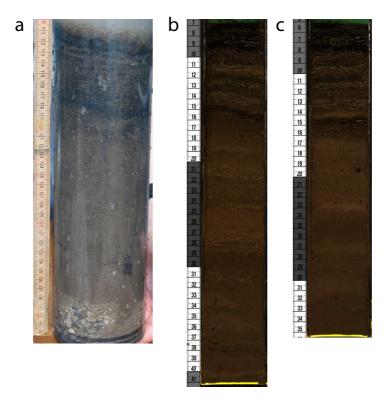


Figure 5.2: a Sediment-surface core CH16-34G, collected in November 2016 at 53-m water depth from near the base of the near-vertical rock face. Sediment-surface cores CH16-14G (b) and CH16-15G (c) were collected in the same month at nearby locations, at 65-m and 60-m water depth, respectively

hence pre-treatment of sediment samples was done following Vaasma (2008). 2359 Four gram of wet sediment was treated for 3 h with 10 mL of 10% HCl at 80°C 2360 in order to remove CaCO₃. Samples were subsequently washed by centrifuging 2361 three times at 3500 rpm, after which 10 mL of H_2O_2 was added to the sediment. 2362 Samples were kept at 80° C and more H_2O_2 was added if necessary until all 2363 organic material was removed (after ~ 2 weeks), and then the suspension was 2364 centrifuged three times at 3500 rpm. In a final step, treatment with 10 mL of 2365 10% KOH for 6 h at 80°C removed (most of the) biogenic silica. After centrifu-2366 gation, samples were stored in Graham's salt (NaPO₃) to avoid coagulation 2367 until execution of grain-size analysis by laser diffraction. 2368

Grain-size measurements were performed using a Malvern Mastersizer 3000. Samples were sonicated for 10 min prior to measurement, after which the suspension was pipetted onto a sieve with 1-mm mesh size and placed into the Mastersizer, until obscuration level reached 7-15%. Three consecutive measurements were performed on the same sample for 12 seconds under continuous stirring at 2500 rpm, and a grain size distribution was generated for each measurement. Additionally, 3-4 g of wet sediment was sieved over a 1-mm mesh-size sieve, and the number of inorganic (either siliciclastic or carbonate) particles >1 mm was counted under a binocular microscope.

2378 5.3.3 Collection of fossil material

The distribution of fossil remains from fish and other lacustrine biota with 2379 depth in the core was assessed by sampling $\sim 2 \text{ cc}$ of wet sediment from each of 2380 the 19 intervals, stirring it in H_2O for 30 min and sieving the resulting solution 2381 into two size fractions (> 250 μ m and 150-250 μ m). These samples were then 2382 stored in Graham's salt (NaPO₃) to prevent coagulation, and counted under a 2383 binocular microscope. Counts included fish fossils (scales, bones), gastropod 2384 molluscs (snails and limpets), ostracods (seed shrimps), cladocerans (water 2385 fleas) and their resting eggs (ephippia), chironomid larvae, sponge spicules. 2386 and non-diagnostic insect and plant remains. 2387

To obtain sufficient fossil fish teeth for numerical analysis, an additional 2388 ~ 20 g of wet sediment from each interval was stirred in H₂O for 30 min. This 2389 solution was then sieved into two size fractions (> 250 μ m and 150-250 μ m) 2390 that were both searched for fossil teeth under a binocular microscope. Teeth 2391 were counted and allocated to their respective jaws (oral or pharyngeal). Oral 2392 teeth were classified into the three principal tooth types (unicuspid, bicuspid 2393 or tricuspid). All recovered teeth were subsequently photographed at 60x 2394 magnification for later measurement. 2395

2396 5.3.4 Data analysis

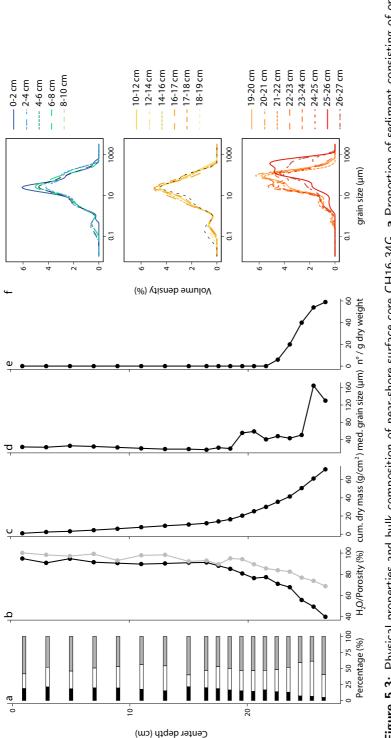
Tooth length was measured using ImageJ v.1.49m (Rasband, 1997) after the
method used in Chapter 2. We used the Malvern Mastersizer 3000 software
v.3.62 to calculate the average grain-size distribution in each depth interval (i.e.
average of three consecutive measurements performed on the same sample).
This average grain-size distribution was then used to calculate a median grain
size for each sediment interval. All further analyses were performed in R

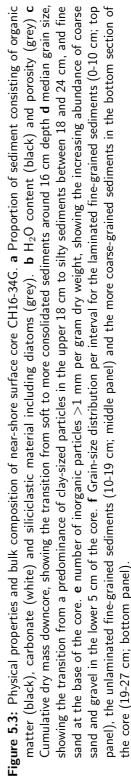
(R Development Core Team, 2016). After assumption testing we examined correlations between fossil abundances and median grain size, and between the median tooth size and median grain size, using Spearman rank tests for non-parametric correlations. Moreover, the proportions of oral and pharyngeal teeth, and of the three oral tooth types, were calculated for each interval to assess changes in tooth assemblage composition between depth intervals, again for comparison with sedimentological variables.

2410 5.4 Results

The upper section of core CH16-34G (0-9 cm depth) consists of very soft, 2411 finely laminated and fine-grained sediments with high water content (>90%); 2412 Fig. 5.3b & 5.3f top panel). Based on the overall thickness of the sequence 2413 of alternating dark and light layers (Fig. 5.2a), this upper section represents 2414 approximately the last 25 years of lake history (i.e. its lower boundary can be 2415 tentatively dated to the early 1990s). The middle section of the core (9-16 cm 2416 depth) consists of similarly soft and fine-grained but unlaminated (massive) 2417 sediments. The organic matter (OM) content of the upper and middle sections 2418 is also similar, and ranges between 15 and 21% (mean 19%; Fig. 5.3a). From 2419 17 cm depth onwards, water content starts to decrease (Fig. 5.3b) but OM 2420 content and median grain size remain stable until 19 cm depth (Fig. 5.3d & 5.3f 2421 middle panel). Median grain size starts to increase below 19 cm, with silt-sized 2422 particles becoming the dominant fraction (Fig. 5.3f bottom panel). Median 2423 grain-size is again slightly lower between 22 and 25 cm (but still silt-sized) and 2424 then rises more dramatically at 25 cm depth with sand-sized particles becoming 2425 predominant at the base of the core (Fig. 5.3d & 5.3f), and OM content drops 2426 to 5-7% (Fig. 5.3a). Clastic mineral particles exceeding 1 mm in size are absent 2427 in the upper 22 cm, but become abundant in the bottom intervals (Fig. 5.3e) 2428 so that the base of the core is distinctly gravely in appearance. These large 2429 particles consist mostly of either pyroclastic material (basalt, lava, tuff) or 2430 carbonate. 2431

The depth distribution of fossils of lacustrine biota other than fish throughout core CH16-34G is highly taxon-specific (Fig. 5.4). Arranging the taxa by increasing median depth of occurrence indicates that chitinous remains (Fig. 5.4a-d: ephippia, 0 cm; chironomid larvae, 14-16 cm; other insect remains,

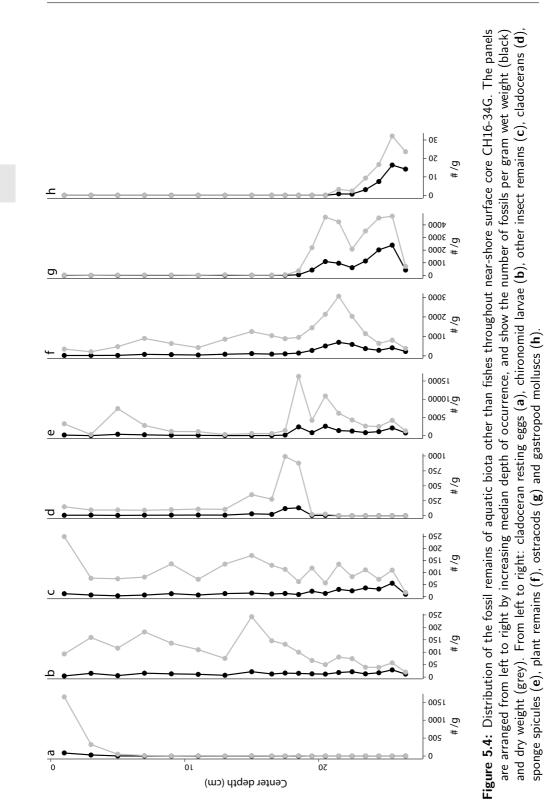




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16-17 cm; cladocerans, 17-18 cm) occur on average higher up in the core than do 2436 fossils consisting of silica (Fig. 5.4e: sponge spicules (Spongilla macroscleres), 2437 19-20 cm) and fossils consisting of carbonate (Fig. 5.4g-h: ostracods, 22-23 cm; 2438 gastropod molluscs, 25-26 cm). However, the relative size of the fossil remains 2439 also matters, with all chitinous fossils being relatively small (<1 mm), and the 2440 carbonate fossils typically larger (2-10 mm). Spongilla macroscleres are small 2441 $(250-350 \ \mu m)$ but consist of solid silica, and assume an intermediate position 2442 with significant presence both in the upper laminated section and in the silty 2443 basal section (Fig. 5.3e). Also, the relatively light-weight but mostly large 2444 terrestrial plant remains are retrieved in high abundance from all sections of 2445 the core, but with a distinct peak in the upper part of the silty basal section 2446 (Fig. 5.4f; median depth 21-22 cm). 2447

Fish fossils are relatively scarce in the fine-grained upper and middle 2448 sections of the core (Fig. 5.5a-c), although a short-lived increase is evident 2449 in the 8-10 cm interval, more prominently so in the abundance of scales and 2450 bones than of teeth. The abundances of all three types of fish fossils increase 2451 more dramatically from 19 cm depth onwards, with a general dip in abundance 2452 between 23 and 25 cm depth. Overall these abundances are significantly 2453 correlated with median grain size (teeth: r = 0.69, p = 0.01; scales: r = 0.56, 2454 p = 0.01; bones: r = 0.68, p = 0.01). The median depth of occurrence of fish 2455 scales (Fig. 5.5a; 19-20 cm) is rather low, being heavily influenced by the peak 2456 in the 8-10 cm interval. The median depths of fish bones (Fig. 5.5b; 23-24 cm) 2457 and fish teeth (Fig. 5.5c; 24-25 cm) are similar to those of other fossils with high 2458 carbonate contents (ostracods and molluscs), indicating similar hydrostatic 2459 properties. There is no significant correlation between median tooth size (Fig. 2460 5.5f) and median grain size (r = -0.10, p = 0.68). Further, the absence of a 2461 correlation between the proportion of pharyngeal versus oral teeth (Fig. 5.5d) 2462 and median grain size (r = -0.36, p = 0.14) suggests that the representation 2463 of teeth from the two jaws in the fossil record is not strongly dependent on 2464 sediment texture. However, a positive correlation can be observed between the 2465 proportion of tricuspids (r = 0.57, p = 0.01) and unicuspids (r = 0.66, p < 2466 (0.01) and median grain size (Fig. (5.5e)). 2467



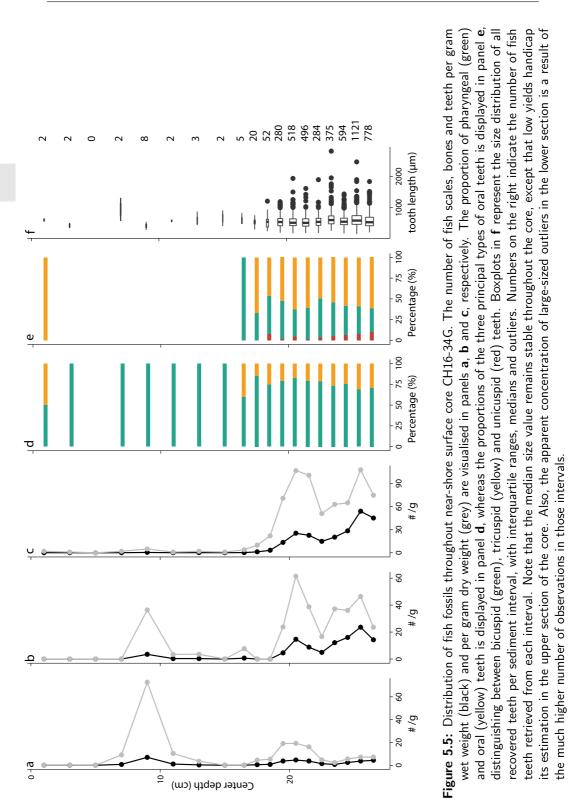
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2468 5.5 Discussion

The large inorganic particles found in the bottom intervals of core CH16-34G 2469 suggest that the gravity corer touched down on rock bottom, and that the 2470 recovered section represents the entire sedimentary deposit present on top of 2471 these rocks at 53 m depth. Alternatively, the nearby 1950s construction of 2472 the pylon for a mechanical lake-level gauge may have caused small pieces of 2473 concrete, likely manufactured with locally available sand and thus difficult to 2474 distinguish from natural rock, to be deposited on top of the soft sediments near 2475 the lake periphery. The high water content of these sediments may have allowed 2476 such concrete fragments to sink through the core, but no evidence of drop 2477 stones traveling through the sediment was found during the extrusion process. 2478 Moreover, several intervals with low water content lie on top of the intervals 2479 with high abundances of large particles, arguing against easy post-depositional 2480 protrusion. 2481

The total amount of dry sediment deposited at this location in the last ~ 25 2482 years equals the value of cumulative dry weight at 9 cm depth (0.78 g/cm^2) . 2483 Assuming that sediment has accumulated at a similar rate before that time, 2484 we can assign an approximate age of 55 years (ca. 1960 AD) to the base of 2485 the fine-grained unit at 19 cm depth. Since there is no evidence that lake level 2486 was significantly lower than today during most of the 20^{th} century prior to the 2487 1990s, we propose that the lack of fine sediment lamination beyond 9 cm depth 2488 is presumably the result of multiple events of unusually deep seasonal-mixing, 2489 the last of which occurred ~ 25 years ago. These events either created enough 2490 bottom turbulence to reshuffle the surficial sediment and destroy its lamination, 2491 or it injected sufficient oxygen to allow temporary development of a zoobenthos 2492 community which would have destroyed the lamination through burrowing (i.e. 2493 bioturbation). Considering that this part of the lake bottom is located >10 m 2494 lower than the usual depth of seasonal oxygen injection today, and that the 2495 oxygen values required for development of a zoobenthos community ($\sim 2 \text{ mg/l}$) 2496 typically persist for only about two months (De Crop *et al.*, unpublished data), 2497 the former mechanism of sediment disturbance may be the most likely. 2498

During such an event, the very soft surface sediment (containing >90%water) is first re-suspended into the lower water column, then resettles with a rate proportional to the size of the particles it is composed of. Large-



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sized sediment grains and biological fossils with high specific density settle 2502 down rapidly after disturbance, whereas re-deposition of smaller-sized mineral 2503 particles (clay and fine silt) and light-weight fossils can take hours to days 2504 (Ferrara & Hildick-Smith, 1982). While in suspension, these small and/or 2505 light particles have a higher probability than the larger and/or heavy particles 2506 (and fossils) to be transported offshore and to eventually settle in deeper 2507 water. Therefore, due to this winnowing process the proportion of large/heavy 2508 particles can be expected to increase with depth in the sediment. This seems 2509 to be the case here, with the relatively abrupt increase in median grain size 2510 between 19 and 21 cm potentially indicating the maximum depth of sediment 2511 disturbance which has occurred during these extreme mixing events since the 2512 1960s. Two important caveats must be mentioned, however. First, since a 2513 substantial fraction of fine-grained sediment may have been winnowed away 2514 from this location during each of these disturbance events, the base of the 2515 middle core section at 19 cm depth may well be significantly older than the 2516 second half of the 20^{th} century. Second, for most of the time which has elapsed 2517 since then, the package of sediment deposited at this location was substantially 2518 thinner than it is today. Thus, even the most severe sediment-mixing event 2519 will not have mixed much more than 10 cm of soft surface muds. Finally, the 2520 decline in median grain size between 22 and 25 cm depth, followed by a steep 2521 increase towards the gravelly base of the core, may be a relic from a previous 2522 phase in lake history with slightly different bottom dynamics in this near-shore 2523 area. 2524

A sequence of short-lived disturbance events is also reflected in the or-2525 dered fashion in which biological fossils are distributed throughout the core 2526 (Fig. 4.4). The hydrostatic properties of large, carbonate-rich fossils such 2527 as gastropod molluscs imply that they are not easily relocated in the first 2528 place, and immediately sink after disturbance. Other calcified and heavy fossils 2529 like ostracods and fish fossils would likewise sink rapidly, followed by sponge 2530 spicules consisting of dense biogenic silica, and only then chitinous fossils such 2531 as cladocerans, chironomid larvae and insect remains. The lightest of all fossils 2532 are the cladoceran ephippia, which consist of a porous chitinous structure. 2533 Their complete disappearance below 3 cm core depth does not necessarily 2534 imply poor preservation, since these resting eggs are commonly preserved in 2535 lake sediments for multiple millennia (e.g., Iglesias *et al.*, 2016). Rather, we 2536

surmise that a large fraction of freshly deposited ephippia is being winnowed away from this near-shore site by bottom currents that are too weak to even disturb the seasonal sediment lamination. The distribution of the densest fossil remains mirrors the double peak observed in median grain size, corroborating our suggestion of the possibility of a second, more ancient series of disturbance events occurring under different conditions of near-shore sedimentation.

The peak in fish fossils observed at 8-10 cm depth likely reflects the 2543 localised disintegration and partial removal of a fish carcass. Overall, however, 2544 the numbers of fish fossils, especially teeth, in the top 18 cm of the core are very 2545 low, suggesting that the slope at this near-shore coring location may expose 2546 also fish fossils to at least some gravity-driven focussing offshore. The majority 2547 of the fossil teeth encountered in this top section are pharyngeal teeth, yet 2548 relative abundances of oral and pharyngeal teeth resemble those encountered 2549 at the lake's depositional centre (Chapter 6). Our results indicate that fish 2550 fossils and the densest fossils of other lacustrine biota possess highly similar 2551 hydrostatic properties, hence the likelihood that fish teeth and bones are being 2552 transported over very long distances, even during extreme deep-mixing events, 2553 must be rather low. This argues against frequent burial at the lake centre of 2554 teeth that initially accumulated elsewhere. We therefore surmise that most 2555 teeth retrieved from offshore core locations, including our 25,000-year sediment 2556 record (Chapter 6) were deposited in situ when fish occasionally ventured 2557 offshore, likely through the process of tooth replacement. Fish carcasses appear 2558 to become permanently buried in articulated form (Verschuren et al., 2017), 2559 as can be expected in the quiet environment of permanently anoxic bottom 2560 waters. In core intervals they are easily distinguished by the strongly increased 2561 local abundance of fish fossils. Post mortem transportation of fish carcasses 2562 has likely occurred (Elder & Smith, 1988), but only during a relatively short 2563 amount of time. Therefore, both individual teeth and fish carcasses preserved 2564 in the sediment represent the lakes' ichthyofauna at the time of their burial. 2565

Our observations in this study thus indicate that sediment focussing currently has very little influence on fish fossil distributions at the lake's depositional centre. This result supports the argumentation that the modest changes in sediment focussing patterns induced by lake-level fluctuations in the last 25,000 years (Moernaut *et al.*, 2010) probably had no influence on fossil distributions offshore. The fractional abundance of teeth derived from

different jaws (i.e. oral and pharyngeal) is not correlated to grain size, but oral 2572 tooth type proportions are to some extent, indicating that severe changes in 2573 sedimentation patterns potentially influence fossil assemblage composition. The 2574 extreme downcore gradient in the texture (grain size) of sediments deposited 2575 near the lake periphery is however never observed in cores from the depositional 2576 centre of Lake Chala, arguing against a strong influence of sedimentology and 2577 taphonomy on its fish fossil assemblages. Overall, fossil assemblages retrieved 2578 from offshore cores likely represent the fish population inhabiting Lake Chala 2579 at that time well. 2580

CHAPTER 6

Long-term trends in oral tooth morphology of *Oreochromis hunteri* linked to past lake-level fluctuations

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2600 6.1 Abstract

Climate-induced habitat changes have often been invoked as important pacers 2601 of evolution in African cichlid fishes, yet studies directly linking morphological 2602 changes to environmental fluctuations are scarce. Here, we track changes 2603 through time in the oral tooth morphology of a cichlid fish in relation to past 2604 lake-level fluctuations. We analysed the variability in oral tooth morphology of 2605 modern-day Oreochromis hunteri, the only indigenous fish species inhabiting 2606 crater lake Chala (Kenya/Tanzania) as a reference for studies of fossil teeth. 2607 Fossils were recovered from six time windows representing alternating high-2608 and low lake-level stands during the last 25,000 years. We find that, as in other 2609 cichlids, modern-day O. hunteri gradually replace their bicuspid/tricuspid oral 2610 dentition with more unicuspid teeth as they grow. Also, O. hunteri has likely 2611 been the only fish species inhabiting Lake Chala throughout the last 25,000 2612 years. However, we observe a recurrent trend linked to lake level in the relative 2613 abundance of oral tooth types, with an increasing proportion of unicuspids 2614 during lowstand periods. This increase is not associated with a larger median 2615 body size, as teeth from lowstands are not significantly larger. Thus, fishes 2616 from lowstand populations developed a unicuspid dentition at smaller size than 2617 today. These shifts are systematic and can be linked directly to climate-driven 2618 habitat change. 2619

2620 6.2 Introduction

The ability of species to adapt their ecology and behaviour to a changing 2621 environment is considered a major driving force of speciation. Especially when 2622 environmental changes give rise to new, unexploited habitats (i.e. ecological 2623 opportunity), rapid adaptation to these habitats provides ample opportunity for 2624 populations to diverge and speciate (Schluter, 2000). Such ecological speciation 2625 is famously prominent in African cichlid fishes (Cichlidae), where it has rapidly 2626 produced highly diverse species flocks (Fryer & Iles, 1972; Johnson et al., 1996; 2627 Salzburger et al., 2014). Both major and minor lake-level fluctuations induced 2628 by Quaternary climate change have been invoked as pacers of cichlid radiation, 2629 alternatingly creating and eliminating particular types of freshwater habitats 2630 (Rossiter, 1995; Danley et al., 2012). 2631

Intrinsic biological factors also influence whether cichlids radiate, and 2632 the colonisation of new lakes and habitats does not automatically result in 2633 diversification and radiation (Gante & Salzburger, 2012; Wagner et al., 2012). 2634 One group of cichlids whose members rarely radiate is the paraphyletic species 2635 group formerly referred to as 'tilapia' (Klett & Meyer, 2002; Dunz & Schliewen, 2636 2013; Brawand et al., 2014). In contrast to the more stenotopic haplochromine 2637 cichlids, tilapiines, including the genus Oreochromis, display a less specialised 2638 but highly versatile morphology, making them more tolerant to habitat change 2639 and less prone to extinction (Ribbink, 1990; Klett & Meyer, 2002). One trait 2640 that has enabled cichlids in general to cope with new environments is great 2641 flexibility of their trophic apparatus, which consists of oral and pharyngeal jaws 2642 (Muschick et al., 2011; van Rijssel et al., 2015). Oreochromines are generally 2643 herbivorous mouthbrooders and hence have similar overall trophic morphologies 2644 to accommodate multiple functions (Trewavas, 1983), yet subtle differences 2645 in jaw and tooth morphology permit them to specialise in certain ecological 2646 niches, especially during periods of food scarcity. 2647

Morphological variation among *Oreochromis* populations is often directly 2648 linked to differences in local environment (Ndiwa et al., 2016), but studies of 2649 naturally-induced morphological shifts in response to changing surroundings 2650 are rare, especially over extended periods. In this respect, the study of fossils 2651 has proven effective for some other groups of teleost fish (Purnell et al., 2007; 2652 Bellwood et al., 2014). But for cichlids, the scarcity of continuous fossil 2653 deposits, the typically disarticulated nature of remains and the limited ability 2654 to assign individual fossils to specific lineages has hampered detailed study of 2655 long-term trends in their ecology and evolution (Murray, 2001). Cichlid teeth 2656 and scales preserved in lake sediments hold the best hope for a continuous 2657 record and are increasingly being explored (Reinthal et al., 2011; Muschick 2658 et al., 2018), but because of confounding factors (Chapter 2) detailed studies 2659 of environment-phenotype relations in cichlid species lineages are still lacking. 2660

The isolated crater lake Chala (Kenya/Tanzania) has a sediment record demonstrating long-term continuity of lacustrine conditions (Verschuren *et al.*, 2009; Moernaut *et al.*, 2010), which contain abundant fossil teeth, bones and scales of cichlid fishes throughout at least the last 25,000 years (Fig. 6.1). The lake currently harbours the endemic Chala tilapia (*Oreochromis hunteri* Günther 1889), which is also the only indigenous fish species. Due to a series

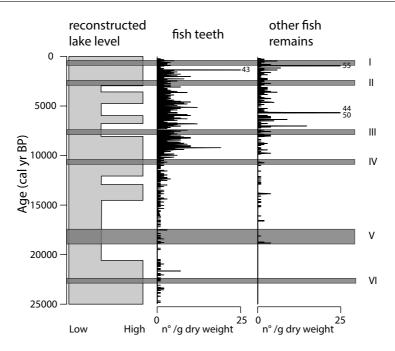


Figure 6.1: Counts of cichlid fossils with respect to reconstructed lake level (after Moernaut *et al.*, 2010). The green bars indicate the six time zones of alternating high-and lowstands that were sampled intensively for fossil teeth.

of anthropogenic introductions in the 1980s, it now shares its habitat with
two other tilapiines, *Oreochromis* cf. *korogwe* and Redbreast tilapia (*Coptodon rendalli* (Boulenger, 1897)), and one haplochromine cichlid (Dadzie *et al.*, 1988;
Seegers *et al.*, 2003).

Lake Chala is the focus of a multifaceted research programme aiming to 2671 reconstruct the long-term environmental and climate history of equatorial East 2672 Africa. These studies revealed that the lake has experienced a sequence of 2673 pronounced lake-level fluctuations during the last 25,000 years (Moernaut et al., 2674 2010; Fig. 6.1). These lake level drops likely had an impact on the nature of 2675 aquatic habitats available to O. hunteri. In Chapter 2, we hypothesised that 2676 morphological change allowing O. hunteri to exploit habitat changes caused 2677 by lake-level fluctuations are reflected in the fossil record, in particular in 2678 the shape of its oral teeth, and that these teeth can be used as a proxy for 2679 changes in trophic ecology. The isolated nature of Lake Chala, its well-studied 2680 environmental past, and the known timing of the arrival of other species 2681 provide a uniquely controlled setting for a detailed morphological analysis of 2682

its fossil record. However, despite being the type species of its genus, only 2683 a general morphological and ecological description of Oreochromis hunteri is 2684 currently available (Trewavas, 1983). Therefore, a quantitative modern-day 2685 reference framework of oral jaw morphology must be constructed before the 2686 fossil record can be optimally explored. In this context, we first provide a 2687 detailed quantification of the intraspecific variability in oral tooth morphology 2688 of living O. hunteri. This is followed by an analysis of the diversity in oral tooth 2689 morphology during six time windows in the last 25,000 years, and assessment 2690 whether any changes in tooth shape observed through time can be linked to 2691 the independently documented environmental change. 2692

²⁶⁹³ 6.3 Material and methods

2694 6.3.1 Study system

The modern-day water level of Lake Chala is relatively high, with a maximum 2695 depth of ~ 90 m (Moernaut *et al.*, 2010). The narrow, rocky shoreline quickly 2696 drops down to a depth of nearly 55 m, from where soft sediments slope more 2697 gently towards the middle of the lake (Fig. 5.1; Opitz, 2008). Monthly 2698 monitoring of dissolved oxygen throughout the water column indicates that the 2699 well-oxygenated zone extends to a maximum of 45 m during mixing events at the 2700 end of the long dry season of July to September. For most of the year oxygenated 2701 water is limited to the uppermost 15 to 30 m (Wolff *et al.*, 2011), thereby 2702 restricting viable fish habitats to the steep rocky crater walls and the open-water 2703 (pelagic) zone. The rocks are covered by epilithic algae and interspersed with 2704 small sandy patches, providing a relatively continuous food source throughout 2705 the year. By contrast, pelagic productivity is largely restricted to a short-2706 lived phytoplankton bloom in August-September (Buckles et al., 2014); low 2707 phytoplankton productivity outside of this period implies poor open-water 2708 feeding conditions during the rest of the year. The depositional centre of Lake 2709 Chala has most probably been continuously anoxic throughout documented 2710 lake history (Verschuren et al., 2009), allowing for excellent preservation of 2711 fossils (Meyer et al., 2018). Nevertheless, several prolonged lake-level drops 2712 in the past 25,000 years were probably severe enough (\sim 30-40 m, Moernaut 2713 et al., 2010) to bring oxygenated conditions to soft-bottom sediments at the 2714 base of the rocks around the lake periphery, which currently occur at depths 2715

Long-term trends in oral tooth morphology of *Oreochromis hunteri* linked to past lake-level fluctuations

too great (at least 55 m; Opitz, 2008) to be supplied with adequate oxygen.
Therefore, soft-bottom benthic habitats may have become available or even
predominant during such lowstands (Fig. 1.6).

2719 6.3.2 Modern-day study material

To study the extant population of O. hunteri in Lake Chala, we obtained 88 2720 fresh-caught specimens purchased from fishermen between January 2014 and 2721 September 2015. Fishermen also determined the sex of these specimens, but 2722 later comparison with photographs of these specimens showed that this sex 2723 information was unreliable (although fishermen proved very adept at species-2724 level identification, Chapter 3). Fish were photographed in lateral view, and 2725 standard length (SL; distance between the tip of the snout and the posterior end 2726 of the last vertebra) was measured to generate a body-size frequency distribution 2727 of the population. The heads of 12 specimens with SL values covering the full 2728 body-size gradient were stored frozen until transport to Belgium. Upon arrival 2729 in the lab they were preserved in 10% formaldehyde for later tooth extraction. 2730 For this purpose, the oral jaws (left dentary, DTL; left premaxilla, PML) were 2731 dissected, cleared and stained using a protocol adapted from Taylor & van Dyke 2732 (1985; for details see Chapter 2). Oreochromis species typically possess very 2733 high numbers of oral teeth implanted in the jaws in multiple rows (Trewavas, 2734 1983; between \sim 350 and 950 in *O. hunteri*; Supplementary Figs. D.3, D.4 & 2735 D.5). We therefore extracted a random subset of 60 teeth from each specimen, 2736 which was determined by systematically lowering the number of teeth included 2737 in the morphometric analysis of two test specimens (Chapter 2), without loss 2738 of information in the generated morphospace (no decrease in convex hull area). 2739 O. hunteri generally has bicuspid outer-series teeth (FR) and tricuspid second-2740 (SR) and inner-series teeth (IR). For this study, we extracted 10 FR, 5 SR 2741 and 15 IR teeth, from both the DTL and PML of each analysed specimen, all 2742 randomly within the respective row classes. 2743

Teeth were photographed at 60x magnification under a binocular microscope. The curved nature of longer teeth hampered standardised orientation of the tooth crown. Therefore, the crowns were clipped off and photographed again in a standardised manner for geometric morphometrics. To optimise morphometrics, the photographs were enhanced in contrast using Photoshop CS6, and PML teeth were mirrored to homologise them with DTL teeth for direct comparison. The edited photographs were bundled per jaw location and per specimen into .tps files using tspUtil version 1.58 (Rohlf, 2013c).

2752 6.3.3 Fossil study material

In 2003 and 2005, the CHALLACEA project recovered a 21.65 m-long sequence 2753 of mostly finely laminated muds from the centre of Lake Chala (Verschuren 2754 et al., 2009). After excision of five turbidites, a 20.82 m-long sequence of 2755 continuous lacustrine sediments represents the last 25,000 years of lake history 2756 (Blaauw et al., 2011). The distribution of fish fossils (teeth, scales, bones) in 2757 Lake Chala sediments was first assessed by recording their presence in each 2758 contiguous 2 cm interval throughout the master composite core sequence below 2759 30 cm depth (a total of 1026 depth intervals, after turbidite excision). This 2760 was done by rinsing 2 g of untreated wet mud through a 250 μ m mesh sieve, 2761 and scanning the retained residue under a binocular microscope at 25-40x 2762 magnification. These counts were visualised using package rioja v.0.9.15 in 2763 R v.3.3.2 (R Development Core Team, 2016). Next, we used a lake-level 2764 reconstruction based on seismic stratigraphy (Moernaut et al., 2010) to select 2765 six time windows of between ~ 500 and 1500-year duration from alternating 2766 lake highstands and lowstands over the last 25,000 years that were certain to 2767 contain sufficient fossil teeth for comprehensive morphometric analyses (Fig. 2768 6.1). Given the considerable sample volumes required to extract substantial 2769 numbers of fossil teeth (383-1858 g of wet mud), the exact length and position 2770 of the selected time windows also depends on the availability of sufficient 2771 core material. The fossil teeth were therefore not sampled from the (already 2772 partly depleted) master core sequence, but from overlapping duplicate core 2773 sections. The visually distinct lamination of Lake Chala sediments permitted 2774 unambiguous cross-correlation of these overlapping sections to the master 2775 sequence. 2776

Extraction of teeth was done by stirring sediments for 30 min in H₂O, after which the solution was sieved into two size fractions (150-250 μ m and >250 μ m) that were both searched for fossils under a binocular microscope. Only oral teeth were retrieved, representing 27% of all fossil teeth encountered. Pharyngeal teeth were counted but not analysed further. The collected teeth were photographed at 60x magnification, clipped re-photographed and treated following the same procedure as adopted on modern-day teeth. Photos of bicuspid teeth that had their major cusp on the left were flipped to ascertain
homology in direct comparisons, but the quasi-symmetrical tricuspid and
unicuspid teeth were left unchanged. Fossil teeth were grouped in .tps files per
time window.

2788 6.3.4 Analysis of tooth size and shape

Analysis of tooth size and shape followed the method described in Chapter 2. 2789 Tooth length and narrowest width were measured in ImageJ v. 1.49m (Rasband, 2790 1997) and further analysed in R. The necks of fossil teeth were often broken, 2791 preventing measurements of their total length, however, the crowns of these 2792 teeth could usually still be included in the morphometric dataset. Therefore, 2793 we used the width of the tooth crown (more specifically, the enameloid width 2794 as represented by the Euclidean distance between landmarks 1 and 3, see 2795 below) as a proxy for tooth length. For this purpose, we first analysed the 2796 correlations between enameloid width (EW), tooth length (TL) and body size 2797 (SL) in modern-day specimens using nonparametric Spearman rank tests, for 2798 each of the three tooth types separately. We then compared the median EW 2799 of fossil teeth from highstands and lowstand episodes using a nonparametric 2800 two-sample Wilcoxon tests, also grouped per tooth type. Nonparametric tests 2801 were used because the assumption of normality for parametric tests was not 2802 fulfilled, a common issue in fish studies because their indeterminate growth 2803 tends to produce skewed size distributions (Nash et al., 2014). Finally, we also 2804 compared median EW of teeth between modern-day and fossil assemblages. 2805

Tooth crown shape was analysed using semi-landmark analysis, in which 2806 two curves containing 20 semilandmarks each were attached to three fixed 2807 landmarks (LM): the left and right lower margin of the enameloid-covered area 2808 of the tooth (LM1 and 3) and the most distal location on the tip (highest cusp) of 2809 the crown relative to the baseline connecting LM 1 and 3 (LM2). The landmarks 2810 were digitised in tspDig2 v.2.17 (Rohlf, 2013a), after which IMP software 2811 was used for Procrustes alignment of landmarks and semilandmarks (Sheets, 2812 2008). Teeth in which unambiguous assignment of one or more landmarks was 2813 problematic (e.g., due to severe abrasion) were excluded, resulting in a dataset 2814 of 671 modern-day teeth and 886 fossil teeth available for further morphometric 2815 analyses (Supplementary Table D.1). All further analyses were performed in 2816 R. To generate a shape space, we ordinated the aligned landmark data of both 2817

modern-day and fossil teeth using non-metric multidimensional scaling in two dimensions (NMDS; packages vegan v.2.4.3 (Oksanen *et al.*, 2017) and MASS v.7.3.45 (Venables & Ripley, 2002)). Ordination of all teeth combined produced an acceptable stress value (8.9) according to the rule of thumb proposed by Kruskal (1964) and Clarke (1993), which indicates that ordination is good if stress values are ≤ 10 . Additional statistical analyses were performed on the NMDS coordinates of each tooth.

The diversity in tooth crown shape present within the modern-day O. 2825 hunteri population was analysed using the NMDS plots of modern-day teeth 2826 only. Although a cichlid's oral teeth are generally described as bicuspid, 2827 tricuspid or unicuspid (e.g., Trewavas, 1983), the assignment of teeth to these 2828 tooth types is not often clear-cut. Hence, to avoid subjectivity, we examined the 2829 robustness of our results under a classification method based on quantitative 2830 morphometry, using angles between cusps. Cusps were considered distinct if 2831 separated by an incision with an angle of less than 135°. Thus, tooth crowns 2832 were classified as tricuspid if they possess two incisions with angles $<135^\circ$, 2833 bicuspids one and unicuspids none. The performance of this classification 2834 method was assessed based on its ability to assign teeth from the modern-day 2835 samples to their known jaw location, and its capacity to reflect changes in 2836 dentition linked to body size. We also compared angle-based classification with 2837 a second quantitative method based on modelled clustering using Gaussian 2838 mixture models (Appendix D). 2839

After a general comparison of fossil tooth shape with modern-day teeth, all 2840 fossil teeth from the two lowstand and four highstand episodes were aggregated 2841 to compare the morphospace occupation of teeth deposited during these two 2842 contrasting habitat phases. We used a permutated Hotelling T^2 -test packages 2843 ICSNP v.1.1.0 (Nordhausen et al., 2015) and Hotelling v.1.0.4 (Curran, 2017) 2844 in two dimensions (NMDS1 and NMDS2) to examine differences, as assump-2845 tions of multivariate normality and common variance-covariance matrices for 2846 parametric tests were not met. This was followed by pairwise comparison of 2847 tooth assemblages from each of the six time windows, likewise using permu-2848 tated Hotelling T^2 -tests. Angle-based classification was then applied to the 2849 ordination data to determine and compare the relative abundances of the three 2850 tooth types. Differences in abundance were validated by randomly resampling 2851 the complete dataset and comparing observed abundances with resampled 2852

²⁸⁵³ abundances of tooth types.

2854 6.4 Results

2855 6.4.1 Modern-day specimens

The standard length (SL) of the 88 O. hunteri specimens collected in Lake 2856 Chala ranged from 16 to 30 cm, with peak distributions (81% of all specimens) 2857 between 23 and 29 cm (Supplementary Fig. 6.2). Tooth length measurements 2858 of modern-day teeth show a clear relation with body size: larger specimens 2859 have longer and wider teeth, and they are more robust (Fig. 6.3). Enameloid 2860 width (EW) correlates strongly with tooth length for all three tooth types 2861 (unicuspids: r = 0.85, p < 0.001; bicuspids: r = 0.82, p < 0.001; tricuspids: r2862 = 0.90, p < 0.001; Fig. 6.4a) and also with body size (unicuspids: r = 0.55, p 2863 < 0.001; bicuspids: r = 0.71, p < 0.001; tricuspids: r = 0.74, p < 0.001; Fig. 2864 6.4b). Therefore, we can use EW as a measure of tooth length for all tooth 2865 types, and fossil EW distributions as an estimate of the median body size of 2866 fossil fish assemblages. 2867

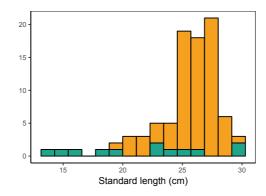


Figure 6.2: Body-size frequency distribution based on standard length of 88 specimens of *O. hunteri.* Colour codes indicate whether teeth were collected (green) or not (yellow).

The shape space containing all 671 teeth from the modern populations is plotted in Figure 6.5a, with the gradient of symbol colour reflecting body size. NMDS-axis 1 discriminates between bicuspids (left, negative side) and tricuspids (right, positive side; Fig. 6.5b), whereas NMDS-axis 2 reflects the

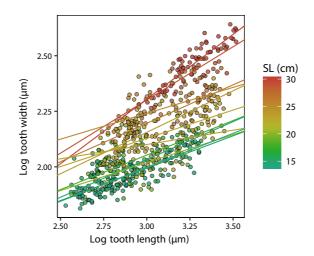
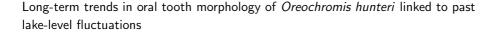


Figure 6.3: Log length-width measurements of modern-day teeth, colour-coded by body size (standard length, SL). Linear regression lines are displayed for each specimen and colour-coded accordingly.

distinction between unicuspids (bottom, negative side) and teeth with multiple cusps (top, positive side; Fig 6.6c). Tooth shape trends related to body size are evident along both NMDS-axes, with teeth from small specimens (more positive values) gradually shifting towards more negative values along both axes, representing more unicuspid teeth in larger specimens.

Shape space occupation of the different tooth types as classified by in-2877 tercusp angle (unicuspid, bicuspid and tricuspid) is visualised in Figure 6.6a. 2878 Despite some overlap, teeth assigned to each of the three tooth types mostly 2879 occupy distinct locations in NMDS shape space. Front-series teeth are predomi-2880 nantly bicuspid, whereas the second and inner series consist mostly of tricuspid 2881 teeth (Fig. 6.6b). Unicuspid teeth occur in all jaw regions. The increasing 2882 trend of unicuspid teeth in larger specimens is clearly reflected in the relative 2883 abundances of each tooth type (Fig. 6.6c), mirroring the patterns observed in 2884 NMDS shape space. In contrast to the model-based approach (Supplementary 2885 information), this classification method thus captures biologically relevant 2886 information and is retained for further examination of tooth-type distribution 2887 in the fossil record. 2888



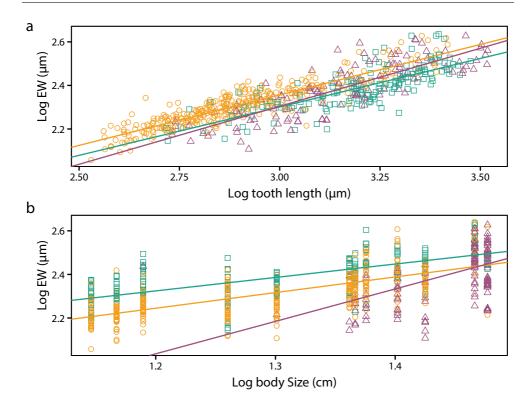


Figure 6.4: Correlations between log enameloid width (EW) and log tooth length (a), and between log EW and log body size (b). Colour codes and symbols represent the different tooth types: unicuspids (purple \triangle), bicuspids (green \Box) and tricuspids (yellow \circ). Linear regression lines are displayed for each tooth type and colour-coded accordingly.

2889 6.4.2 Fossil teeth

Fossil fish remains are (almost) continuously present throughout 25,000-year sediment record of Lake Chala (Fig. 6.1), but their abundances vary substantially through time. The most notable pattern is a marked increase around 10,500 years ago, i.e. at the onset of the Holocene.

Enameloid width (EW) distributions show that the recovered fossil oral teeth (n = 886) are on average significantly smaller than our reference collection of modern-day teeth of *O. hunteri* (Fig. 6.7a; W=114,790; p < 0.001). Aggregating fossil teeth assemblages according to lake level reveals that teeth deposited during two lowstand episodes (subtotal n = 526) are significantly larger, on average, than those from four highstand episodes (subtotal n =

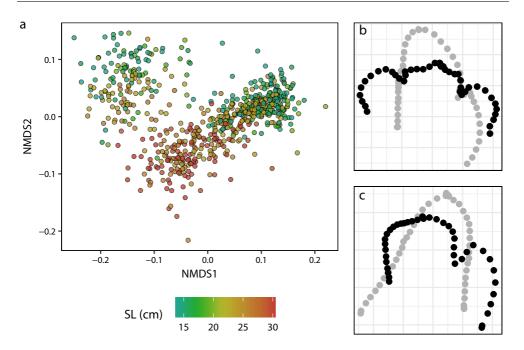


Figure 6.5: NMDS shape space of modern-day teeth with colour reflecting body size (**a**), with outline drawings of specimens representing minimum (grey) and maximum (black) values along NMDS-axis 1 (**b**) and along NMDS-axis 2 (**c**).

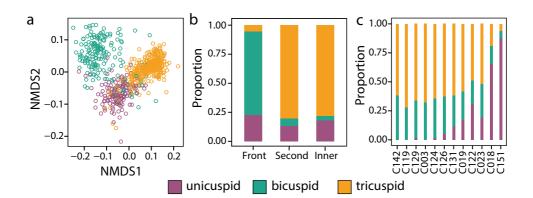


Figure 6.6: a Modern-day NMDS shape space colour-coded by tooth type as classified by intercusp angle, and **b** distribution of tooth types in the jaw. **c** displays proportions of tooth types in each specimen, ordered by increasing body size.

 $_{2900}$ 360; W = 77,434; p < 0.001). However, non-aggregated assemblages display a marked decline in median tooth size from older to younger time windows in all

three tooth types (Fig. 6.8), indicating that the previous result may have been biased by the large sample from lowstand episode III (Supplementary Table D.1), and that the trend in median tooth size is potentially influenced by time passed since deposition, rather than by lake level at the time of deposition. Although most samples are skewed towards larger teeth, this skewness is enhanced in older samples (Fig. 6.8), suggesting that post-depositional corrosion results in a disproportionate loss of smaller teeth over long time periods.

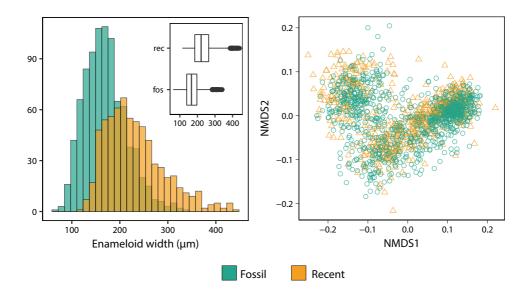


Figure 6.7: a Enameloid width (EW) distributions of fossil (green) and modern-day teeth (yellow), with boxplots (inset) capturing median EW and interquartile ranges. **b** NMDS shape space including both fossil (green) and modern-day (yellow) teeth.

Fossil teeth are uni-, bi- and tricuspid, similar to those in modern-day 2909 O. hunteri. Viewing morphological variation in the oral tooth crowns as a 2910 continuum through time, the variability in tooth shapes represented by our 2911 six fossil assemblages overlaps almost completely with that in the modern-day 2912 population of O. hunteri (Fig. 6.7b), strongly suggesting that all recovered 2913 fossil teeth belong to ancestral O. hunteri. When fossil assemblages are ag-2914 gregated according to lake level, mean oral tooth shape of O. hunteri living 2915 during past lowstand phases is significantly different from that of O. hunteri 2916 during highstand episodes ($T^2 = 85.25$, p < 0.001; Fig. 6.9a). The difference 2917 is mainly situated along NMDS2, with on average more negative values for 2918 'lowstand' teeth. The variation in tooth shape explained by the two NMDS-axes 2919

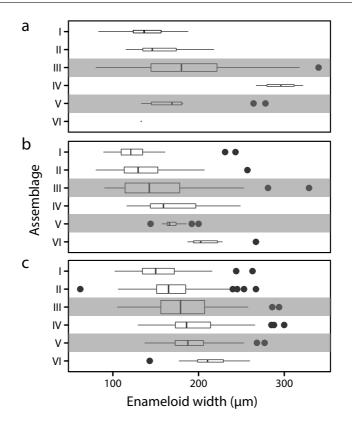


Figure 6.8: Boxplots representing median EW values and interquartile ranges per fossil assemblage, and per tooth type: **a** unicuspids, **b** bicuspids and **c** tricuspids. Grey bars indicate fossil assemblages derived from lowstand periods

is identical to that in modern-day teeth (Fig. 6.5b-c), with more negative values 2920 along NMDS-axis 2 similarly indicating a greater proportion of unicuspid teeth. 2921 Moreover, when comparing the two lowstand and four highstand assemblages 2922 separately, this shift in tooth morphology is shown to be recurrent through time 2923 (Fig. 6.9b). More specifically, permutated Hotelling T^2 -tests yield significant 2924 differences in five out of eight Bonferroni-corrected pairwise comparisons be-2925 tween NMDS-axes values of individual high- and lowstand assemblages (Table 2926 6.1).2927

Plotting fossil oral teeth according to tooth type as classified by intercusp angle reveals a higher amount of overlap between types than in modern-day shape space (Fig. 6.10a). Nevertheless, consistent with our observations in NMDS shape space, the proportion of unicuspid teeth as classified by intercusp Long-term trends in oral tooth morphology of *Oreochromis hunteri* linked to past lake-level fluctuations

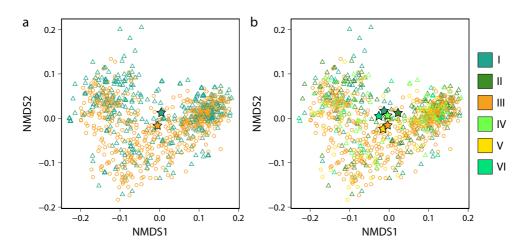


Figure 6.9: NMDS shape space of fossil teeth displaying the distinction between fossil crown shape during lowstands (yellow circles) and highstands (green triangles; a), and between fossil crown shape of six assemblages from alternating high- and low lake level phases (b). Stars indicate centroid location.

Table 6.1: Bonferroni-corrected p-values of permutated Hotelling T²-tests for each pairwise comparison between assemblages. Comparisons between assemblages from contrasting lake-level stands have grey cell colours. P-values significant on the 5% signifance level are indicated with an asterisk (*).

		H I	H II	L III	H IV	L V
Н	II	0.74				
\mathbf{L}	III	$< 0.01^{*}$	$< 0.01^{*}$			
Η	IV	1.00	1.00	0.03^{*}		
\mathbf{L}	\mathbf{V}	0.02^{*}	$< 0.01^{*}$	1.00	0.17	
Η	VI	1.00	1.00	1.00	1.00	1.00

²⁹³² angle is markedly higher during lowstand episodes than during the highstand ²⁹³³ episodes before and after (Fig. 6.10b). The average increase ($\sim 10\%$) is ²⁹³⁴ significantly (p < 0.01) higher than what can be expected by chance.

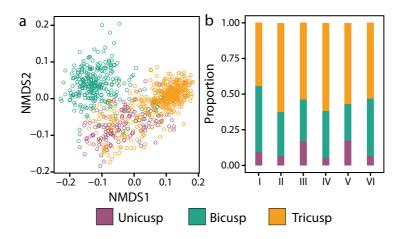


Figure 6.10: NMDS shape space of tooth types assigned using the intercusp angle (a), and their relative proportions in the six sampled time intervals (b).

²⁹³⁵ 6.5 Discussion

2936 6.5.1 The oral dentition of modern-day O. hunteri

Our data show that during their lifetime, modern-day Oreochromis hunteri in 2937 Lake Chala undergo significant shifts in oral dentition. First, the teeth become 2938 gradually wider and more robust with increasing body size (Fig. 6.3). Secondly, 2939 larger fishes develop a greater proportion of unicuspid teeth (Figs. 6.5 & 6.6). 2940 Specifically, small fish have bicuspid FR teeth followed by tricuspid SR and 2941 IR teeth, whereas very large fishes possess an almost entirely unicuspid oral 2942 dentition. Individuals of intermediate body size, which make up the majority 2943 of the specimens available for this study, have an intermediate dentition: they 2944 largely retain the bicuspid/tricuspid arrangement, but the major/middle tooth 2945 cusp becomes more dominant relative to its side cusp(s). The shift towards 2946 a predominantly unicuspid dentition in very large specimens is thus gradual 2947 rather than abrupt (Fig. 6.6). 2948

The shift to unicuspid FR teeth or an entirely unicuspid oral dentition is known to occur in large individuals of several other *Oreochomis* species, mainly in sexually mature males (Trewavas, 1983). Most cichlids display dietary shifts as they grow (Njiru *et al.*, 2008), and these are regularly accompanied by changes in different aspects of their trophic morphology (Streelman *et al.*, 2007). To our knowledge, this is the first study documenting such ontogenetic

change in oral dentition with geometric morphometrics, and it reveals that, at 2955 least in O. hunteri, the changes ensue gradually until completely transforming 2956 the dentition. Oreochromis species are generally herbivorous (Trewavas, 1983), 2957 including modern-day O. hunteri in Lake Chala (Chapter 4). Given the lack of 2958 submerged macrophytes and a reliable pelagic food source, it feeds mostly on 2959 epilithic algae growing on the steep rocky shores (Chapter 4). In other cichlid 2960 taxa a unicuspid oral dentition is often linked to feeding at a higher trophic 2961 level (Fryer & Iles, 1972; Witte & Van Oijen, 1990), but oreochromines are 2962 rarely entirely piscivorous or insectivorous. We therefore surmise that the shift 2963 to unicuspid dentition in sexually mature O. hunteri is associated with their 2964 enhanced territorial behaviour linked to breeding (Trewavas, 1983). 2965

²⁹⁶⁶ 6.5.2 Trends in fossil teeth abundance and preservation

The presence of fossils throughout the core demonstrates that fish have been 2967 (almost) continuously present in Lake Chala over at least the past 25,000 years 2968 (Fig. 6.1). Lack of a systematic contrast in the abundance of fish fossils between 2969 highstand and lowstand episodes argues for a negligible influence of changes 2970 in *post-mortem* transport on the abundance of fish fossils recovered from the 2971 mid-lake sediment record (Chapter 5). Nevertheless, fossil abundances vary 2972 considerably during this period, showing a distinct rise at the Pleistocene-2973 Holocene transition approximately 10,500 years BP. The organic-carbon content 2974 in Lake Chala sediments shows a marked increase in aquatic primary production 2975 during this transition (Blaauw et al., 2011; Barker et al., 2013; Meyer et al., 2976 2018), which is presumably associated with regional climate warming around 2977 the same time (Loomis et al., 2017) and may have enhanced the lake's carrying 2978 capacity to sustain a larger fish population. Based on genetic analyses, we 2979 found no evidence of a pronounced demographic expansion within the last 2980 25,000 years (Chapter 3), although the sample size was relatively small for the 2981 used approach. Taphonomic effects cannot be excluded at this stage. Chemical 2982 corrosion makes older teeth more brittle and possibly reduces their preservation 2983 potential, which not only decreases the overall abundance of fossils in older 2984 samples, but also the proportion of small teeth. However, the relatively modest 2985 trend in mean tooth size with age (Fig. 6.8a) compared to the abrupt, order-2986 of-magnitude increase in fossil abundance at the start of the Holocene (Fig. 2987 6.1) suggests that the latter mostly reflects a true increase of the Lake Chala 2988

2989 O. hunteri population at that time.

Our data on enameloid width (EW) indicates that fossil teeth are on 2990 average smaller than those of modern-day fishes (Fig. 6.7a). No specimens 2991 smaller than 16 cm were collected in the modern-day survey, presumably due to 2992 large fish-net mesh. Fry and juvenile fishes are abundantly present in natural 2993 populations, and typically experience high mortality (Cushing, 1974; Peterson 2994 & Wroblewski, 1984). Teeth of these small fishes should be well-represented 2995 in the fossil record, except that fry teeth may suffer more chemical corrosion 2996 affecting their preservation (and in any case are not retained by the used sieve 2997 if smaller than 150 μ m). Moreover, cichlids continuously replace their teeth 2998 as they grow (polyphyodonthy; Fryer & Iles, 1972; Hulsey et al., 2016), and 2999 a tooth has an average lifespan of around 100 days (Tuisku & Hildebrand. 3000 1994). If O. hunteri has a life span similar to other Oreochromis species (up 3001 to ~ 10 years; Froese & Pauly, 2017), each large specimen has contributed 3002 up to three dozen of generations of teeth to the sediment record, increasing 3003 the proportion of small teeth further. The smaller median size of fossil teeth 3004 thus conforms to expectation. In effect, the minimum size of teeth in our 3005 modern-day reference collection is controlled by the mesh size of the fish nets, 3006 whereas the minimum size of teeth in the fossil assemblage is controlled by the 3007 sieve mesh used for sediment processing. Recent and fossil assemblages are 3008 hence not directly comparable in terms of tooth size, but their tooth shapes 3009 show great similarities (Fig. 6.7b). This implies that fossil O. hunteri tooth 3010 morphology can be interpreted within the modern-day framework. 3011

3012 6.5.3 Temporal trends in oral tooth shape

The large overlap between fossil and modern-day oral teeth in shape spaces 3013 (Fig. 6.7b) confirms that O. hunteri has been the only fish species inhabiting 3014 Lake Chala before the recent introductions, and this for at least the last 25,000 3015 years. The overall variability in tooth shape existing in the current population 3016 has thus been remarkably stable over time. However, we observe significant 3017 and systematic trends in occupied morphospace that can be related to the 3018 documented changes in lake level (Figs. 6.9 & 6.10). During periods of low 3019 lake level, the dentition shifts towards more unicuspid, and we found this trend 3020 to occur repeatedly over time. The observed trends are clearly reflected in the 3021 relative abundances of different tooth types, in which an average increase of 3022

10% in the abundance of unicuspid teeth is observed during lowstand periods(Fig. 6.10).

The morphological changes from highstands to lowstands are overall similar 3025 to the changes in oral dentition occurring in the modern-day population as 3026 specimens grow larger. However, fossil teeth from lowstand assemblages are not 3027 systematically larger than those from highstand assemblages (Fig. 6.8), hence 3028 a larger median body size during low lake levels is not a plausible explanation 3029 for the observed shifts in tooth shape. Rather, our data suggest that, during 3030 low lake level phases, O. hunteri shifted towards a unicuspid dentition earlier 3031 in life, likely triggered by a change in available habitat. In cichlids, unicuspids 3032 are often related to feeding at a higher trophic level (Fryer & Iles, 1972), and 3033 in Lake Chala may have been advantageous to exploit a new food source that 3034 is present in the soft sediments made accessible during lowstands, such as 3035 insect larvae, annelid worms and other macrobenthos. However, insectivorous 3036 oreochromines are rare, although juveniles may opportunistically feed on insects 3037 or even fish fry (Njiru et al., 2004). Alternatively, an earlier shift towards 3038 unicuspid teeth might indicate that sexual maturity is reached at smaller body 3039 sizes (i.e. at an earlier age) during low lake level phases than during highstands. 3040 Intraspecific population variability in body size and age at maturity is common 3041 in Oreochromis species (Duponchelle & Panfili, 1998) and is strongly linked to 3042 environmental factors, such as temperature and population density (Brummett, 3043 1995). 3044

Both phenotypic plasticity (Bouton et al., 2002; Muschick et al., 2011; van 3045 Rijssel et al., 2015) and genetic differences (Albertson & Kocher, 2006; Hulsey 3046 et al., 2017) have been invoked to explain rapid morphological change in the 3047 trophic ecology of cichlid fishes. We can currently only hypothesise on the 3048 exact ecological and evolutionary mechanisms underlying the observed trends 3049 in oral tooth shape, yet the observation in itself is remarkable. Recurrent shifts 3050 in oral tooth shape directly linked to lake-level fluctuations provide long-term 3051 evidence of the ability of *O. hunteri* to respond to climate-driven environmental 3052 variation, and confirm the versatility of this species over extended periods of 3053 time. Rapid morphological responses in the trophic apparatus of O. hunteri 3054 may thus have enabled this species to cope with past environmental change. 3055

CHAPTER 7

General discussion

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7.1 Producing a long-term record of evolutionary change in cichlid fishes

For almost a century, cichlid fishes populating freshwater bodies in both 3062 the Old and New World tropics have been under study as a model system 3063 for ecological speciation and radiation (Fryer & Iles, 1972; Stiassny & Meyer, 3064 1999). Although this research has led to a better understanding of the processes 3065 governing evolution, many questions as to how various factors contributed 3066 to differentiation and reproductive isolation remain unanswered. A major 3067 drawback is the lack of a continuous fossil record of this group (Murray, 2001), 3068 hampering long-term inquiries of morphological patterns and their underlying 3069 processes. Indeed, rapid morphological responses to environmental change, 3070 especially of ecologically important traits, have been documented to occur on 3071 very short timescales (e.g., Witte et al., 2008; van Rijssel et al., 2015) and they 3072 have been proposed to function as a driver of diversification and speciation 3073 in cichlids (e.g., Meyer, 1990; Muschick et al., 2011; Gunter et al., 2013). 3074 Extending morphological observations beyond the historical time scale could 3075 provide long-term evidence of such morphological responses and corroborate 3076 their importance in cichlid evolution, yet these studies are currently lacking. 3077

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Fossil remains of cichlid fishes, like teeth, scales and bones are, however, 3078 often well-preserved in lake sediments. These sediments may in some lakes 3079 provide a continuous palaeontological record, stretching back to the lake's early 3080 origins. Cichlid fossils have been tentatively explored in the sediment record of 3081 the great African rift lake Malawi (Reinthal et al., 2011), but the enormous 3082 species diversity and major past lake-level fluctuations in this lakes (Chapter 3083 1) complicate the analysis and interpretation of its fossil records. Appointing 3084 fossils to a certain species is difficult in these lakes, and changes in the position 3085 of the shoreline likely influenced the composition of fossil assemblages at any 3086 fixed coring location. Ideally, a small and isolated system with limited species 3087 diversity yet a continuous fossil record is examined first to explore the feasibility 3088 of paleontological studies of cichlids as a baseline for later investigations in 3089 more complex systems. Moreover, if this small system has a well-resolved 3090 climatic past and has been prone to environmental fluctuations over extended 3091 periods of time, it could be used to document morphological change in response 3092 to such fluctuations and it may provide insight into long-term ecological and 3093 evolutionary processes underlying cichlid diversity. 3094

This thesis had the general aim of tracing long-term morphological change 3095 in a single cichlid species in response to past lake-level fluctuations. The 3096 isolated crater lake Chala (Kenya/Tanzania) harbours only one indigenous fish 3097 species, the cichlid Oreochromis hunteri Günther, 1889. The lake has been 3098 prone to major lake-level changes throughout at least the last 25,000 years 3099 (Moernaut et al., 2010), which presumably impacted habitat diversity and 3100 niche availability for O. hunteri (Chapter 1). Lake Chala's finely-laminated 3101 and continuous sediment record also contains abundant cichlid fossils (Fig. 3102 1.5), of which especially the morphology of fossil teeth may reflect the species' 3103 past ecology and could therefore be used to trace how feeding habits changed 3104 in function of environmental changes. Lake Chala thus provides a controlled 3105 setting that is ideal to develop studies of long-term morphological change in 3106 cichlids in response to climatically-driven hydrological fluctuations in the past. 3107

Chapter 6 presents the findings as to the overall aim to document longterm patterns in oral tooth shape of *Oreochromis hunteri* in response to strong lake-level fluctuations. As this study is the first of its kind in cichlid fishes, it required a feasibility analysis and methodological exploration, which I developed in Chapter 2 and further improved in Chapter 6. Moreover, to fully understand

morphological trends observed in the fossil record, an integrated approach 3113 that included the extant ichthyofauna as a modern-day reference framework 3114 was required. As Lake Chala fishes had thus far been relatively unexplored, 3115 I investigated species integrity, colonisation history and feeding habits of the 3116 enigmatic O. hunteri in Chapter 3 and 4, respectively, with special attention 3117 to the impact of anthropogenically introduced cichlid species on O. hunteri. 3118 Chapter 6 also includes a detailed description of oral tooth shape in extant O. 3119 hunteri. Beyond these data on the extant fishes, insight into the taphonomic 3120 processes that affect how fossil teeth accumulate in the fossil record is required. 3121 Therefore, the impact of sedimentation dynamics and taphonomic processes 3122 on the distribution of fossil teeth in Lake Chala sediments was assessed in 3123 Chapter 5. In the following discussion, I aim to integrate the findings of these 3124 respective Chapters. I start with a thorough assessment of the methods used, 3125 and elaborate on their potential application in future studies of other systems. 3126 I then discuss the results of my examination of the extant ichthyofauna of Lake 3127 Chala, touching on species integrity and conservation of O. hunteri in light of 3128 recent anthropogenic introductions and integrating the obtained insights in its 3129 diet and oral tooth shape. Important taphonomic influences on the distribution 3130 of tooth fossils in the sediment are discussed before I elaborate on the origin of 3131 Oreochromis hunteri in Lake Chala and its eco-morphological response to past 3132 environmental change. 3133

3134 7.2 Quantitative analysis of tooth shape

3135 7.2.1 Geometric morphometrics for cichlids

The use of geometric-morphometrics to quantify organismal shape is now 3136 well-established in many fields of biology (Bookstein, 1991; Zelditch et al., 3137 2004), and a large body of literature exists of its application to the study of 3138 cichlid fishes (reviewed in Kerschbaumer & Sturmbauer, 2011). In this thesis, 3139 geometric-morphometrics were used in two aspects of the research, namely 3140 to first describe the whole-body morphology of Lake Chala Oreochromis, and 3141 then to describe the shape of their extant and fossil oral teeth. Whole-body 3142 morphology is described in Chapter 2 using a landmark-based approach. This 3143 method was capable to consistently distinguish the two *Oreochromis* species 3144 currently inhabiting Lake Chala: the endemic O. hunteri and the introduced 3145

Oreochromis cf. korogwe. In Chapter 3, the same method was found to be effective to distinguish all Oreochromis species currently present in the Upper Pangani River region, even between the phylogenetically close species O.hunteri and O. jipe.

Fossil fish remains in Lake Chala sediments (and in other lakes; Reinthal 3150 et al., 2011) consist largely of individual teeth, yet contemporary ecomorpholog-3151 ical studies on cichlids mostly analyse complete oral (Albertson & Kocher, 2001; 3152 van Rijssel et al., 2015) and pharyngeal jaws (Hellig et al., 2010; Muschick 3153 et al., 2011). Geometric morphometrics have been applied successfully to 3154 individual fossil teeth of other biota, for example hominids (Martinón-Torres 3155 et al., 2006; Gómez-Robles et al., 2007), sharks (Nyberg et al., 2006; Whitenack 3156 & Gottfried, 2010) and teleost fish (Bellwood et al., 2014; Gauchey et al., 3157 2014), but studying individual teeth has rarely been attempted for cichlids 3158 (but see Wautier et al., 2002). I therefore developed a geometric-morphometric 3159 approach tailored specifically to the analysis of oral tooth shape in the Ore-3160 ochromis species of Lake Chala, O. hunteri and O. cf. koroque. A working 3161 method required solutions to several obstacles, first and foremost the problem 3162 that the diversity in tooth shape within a single specimen is generally much 3163 larger than interspecific differences. Indeed, the oral jaws of both species 3164 contain an outer series of bicuspid teeth and several inner series of tricuspid 3165 teeth (Fig. 2.2). Because of this intraspecific disparity, a sensitive method 3166 was required to distinguish between the oral teeth from different species, and 3167 to trace subtle changes in those oral teeth through time. Moreover, fossil 3168 teeth are often broken in ways that do not allow reconstruction of general 3169 tooth morphology but nevertheless have their tops (crowns) intact, potentially 3170 because the strongly mineralised enameloid layer (Sasagawa, 1997) protects 3171 crowns from breaking. I therefore focused on documenting shape variation in 3172 crowns because it allows maximal data retrieval from the fossil record. I found 3173 that semi-landmark analysis (SLM) followed by principal component analysis 3174 (PCA) on the tooth tops best suited the needs. 3175

Semi-landmark analysis was put to the test in Chapter 6, where extension of the analysis to 12 extant *O. hunteri* specimens of different body sizes revealed considerable ontogenetic change in oral tooth shape. The gradual nature of these changes was clearly reflected in the generated morphospace (Fig. 6.5), indicating that SLM could not only separate ontogenetic extremes but also more

subtle shape differences. However, the gradual transition from bicuspid and 3181 tricuspid oral teeth to unicuspid teeth in large individuals resulted sometimes 3182 in uncertainty as to assigning teeth to tooth types. I therefore also compared 3183 the results of two classification mechanisms, one based on angles between 3184 cusps and the other based on all the variation in the morphometric dataset 3185 and Gaussian mixture models. Moreover, I found a strong linear relationship 3186 between tooth size and body size, which was then used for inferences about 3187 body size in the past. 3188

The very high numbers of oral teeth in the premaxillae and dentaries of 3189 *Oreochromis* species required selection of a randomised yet morphologically 3190 informative subset of teeth from each analysed modern-day specimen. This 3191 subsampling strategy was developed in Chapter 6, comprised 60 teeth per 3192 individual and, after exclusion of teeth that could not be digitised, resulted in 3193 a modern-day reference frame consisting of a total of 671 teeth. Yet these high 3194 numbers of oral teeth had to be matched by a comparable amount of fossil 3195 teeth being analysed if subtle shape shifts through time were to be detected. 3196 Eventually, 886 fossil teeth were included in the morphometric dataset of 3197 Chapter 6, but sample sizes varied greatly between different time windows, so 3198 that non-parametric statistical tests had to be applied to analyse the digitised 3199 datasets. Non-metric multidimensional scaling (NMDS) returned highly similar 3200 results as PCA, but it has less assumptions and, if stress values are acceptable, 3201 it allows representing all shape variation on a predefined number of axes (here 3202 two). Overall, the methods presented in Chapter 6, largely based on earlier 3203 findings in Chapter 2, proved ideal to trace oral tooth shape change within a 3204 species through time in relation to past environmental change. 3205

³²⁰⁶ 7.2.2 Potential of the developed method for future studies

The method of tooth shape analysis used in this thesis can be applied on 3207 both extant and fossil cichlid species in other systems. Combining geometric-3208 morphometric analysis of the tooth crown with length-width measurements 3209 of the neck allows distinguishing between species (Chapter 2) and tracing 3210 subtle shape changes associated with ontogeny (Chapter 6). It could therefore 3211 potentially be implemented in morphological studies of tooth development, 3212 for example enabling more detailed analyses of tooth replacement patterns. 3213 Currently, such studies use qualitative tooth shape descriptions (e.g., Huysseune, 3214

1995; Streelman et al., 2003), yet a quantitative approach of oral tooth shape 3215 analysis could facilitate comparison between studies. It could also be used to 3216 study convergence in oral tooth morphology between more distantly related 3217 taxa, as has for example been done for the shape of the lower pharyngeal 3218 element (Muschick et al., 2012). Although this method is very sensitive for 3219 subtle changes in oral tooth crown shape, such studies may nevertheless require 3220 extra biometric measures for thorough description of complete oral tooth shape 3221 and interspecific differences. For example, tooth curvature is an important 3222 source of shape variation in oral cichlid teeth (Fryer & Iles, 1972) that has 3223 been standardised in this analysis. 3224

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The main aim was however the development of a tool that enables a 3225 quantitative analysis of individual fossil cichlid teeth in lake sediments. I 3226 successfully implemented this approach on the fossil record of Lake Chala, yet 3227 some important caveats should be made for other systems. First and foremost. 3228 very high numbers of fossil teeth are required for complete examination of 3229 shape diversity in the fossil record. In total, 5 kg of wet sediment extracted 3230 from cores was sieved and analysed for the results presented in Chapter 6 3231 alone. Many large scientific drilling programs become increasingly multi- and 3232 interdisciplinary (Wilke et al., 2016), restricting the amount of sediment that 3233 can be dedicated to the analysis of a single palaeoenvironmental proxy. A 3234 sampling strategy generating large amounts of sediment is thus required to 3235 adequately study fossil cichlid teeth, and is ideally delineated before coring is 3236 initiated. 3237

The workload associated with a comprehensive analysis of modern-day 3238 tooth shape is also considerable. I found that incorporation of oral teeth from 3239 several individuals of different body sizes is a prerequisite in morphologically 3240 versatile genera like Oreochromis (Chapter 6), which occupy a wide range 3241 of habitats throughout their lives and functionally adapt their trophic traits 3242 accordingly (Trewavas, 1983; Ribbink, 1990). Stenotopic lineages, such as 3243 the haplochromines, have an ecomorphology that is more closely linked to 3244 the exploitation of very narrow niches (Fryer, 1959b; Fryer & Iles, 1972; Rib-3245 bink et al., 1983; Reinthal, 1990), and for such taxa smaller sample sizes may 3246 suffice to obtain a complete image of extant intraspecific diversity in tooth 3247 shape. Changes in oral tooth shape associated with ontogeny have neverthe-3248 less also been documented in haplochromine species (Streelman et al., 2003, 3249

2007). Moreover, the enormous species diversity combined with considerable convergence of trophic structures among haplochromines (Chapter 1) may impede assignment of individual fossil teeth to a certain species or even genus. Especially in the cichlid species flocks of the African Great Lakes, studying general ecotypes rather than species-specific ecomorphology may provide a viable alternative.

³²⁵⁶ 7.3 The extant ichthyofauna of Lake Chala

The presence of only one indigenous cichlid species (*Oreochromis hunteri*) in 3257 Lake Chala made interpretation of shape changes in fossil oral tooth assemblages 3258 relatively straightforward, as all teeth could be assigned to this species (Chapter 3259 6). However, since very little was known about the extant O. hunteri at the 3260 start of this project, development of a comprehensive modern-day framework 3261 was an important prerequisite for thorough understanding of its fossil record. 3262 In this respect, the recent introductions of three other cichlid species by humans 3263 (Oreochromis cf. korogwe, Coptodon rendalli and Haplochromis sp. "Chala") 3264 potentially complicated this study, as interspecific hybridisation with one of 3265 these species could compromise the extant trophic morphology of *Oreochromis* 3266 hunteri (Holzman & Hulsey, 2017). This is especially relevant in Lake Chala, 3267 where a deep open-water environment surrounded by near-vertical crater walls 3268 offers limited potential for reproductive or niche segregation. Moreover, a 3269 correct interpretation of changes in oral tooth shape requires understanding of 3270 the feeding habits of modern-day O. hunteri. 3271

3272 7.3.1 Species integrity of Oreochromis hunteri

To examine the influence of hybridisation, I first studied the genetic and 3273 morphological integrity of Oreochromis hunteri. Hybridisation is most likely to 3274 occur between congeneric species, hence I focussed on potential interbreeding 3275 of O. hunteri with O. cf. korogwe. In Chapter 2, I first performed geometric 3276 morphometrics to assess body and oral tooth morphology of both taxa and 3277 found their overall morphology to be clearly distinct, already arguing against 3278 recent and/or ongoing hybridisation. In Chapter 3, I extended this study 3279 with a molecular-genetic assessment of two loci in the mitochondrial DNA 3280 (mtDNA) of O. hunteri and O. cf. koroque, and found no signs of introgression 3281

between the two. I therefore concluded that both the genetic and morphological integrity of *Oreochromis hunteri* is currently intact, and that it can be used as a modern-day representative for *O. hunteri* remains in the fossil record of Lake Chala.

Potential mechanisms of reproductive isolation acting between O. hunteri 3286 and O. cf. koroque are briefly touched upon in Chapter 3, and include separa-3287 tion in spawning time and distinctions in microhabitat preference (for example 3288 described in Pullin & Lowe-McConnell, 1982; Lowe-McConnell, 1987; Beveridge 3289 & McAndrew, 2000). Although tilapia in near-equatorial regions generally 3290 spawn year-round, many species display breeding peaks at a certain time of the 3291 year (e.g., Lowe-McConnell, 1987; Brummett, 1995). It is not known whether 3292 the two Lake Chala *Oreochromis* species have a delineated spawning time. 3293 If they do, then fishes probably stick to small sandy patches for spawning 3294 and substrate brooding (in Coptodon rendalli), whereas mouthbrooding Ore-3295 ochromis females often reside in sheltered areas (Lowe-McConnell, 1987) and 3296 thus, in Lake Chala, likely hide out in crevices between rocks. Obtaining fish 3297 is generally difficult for all tilapiine species at the end of May and in June, 3298 and in November and December, when nets are lowered and fishermen resort 3299 in part to line fishing to supplement the net catches (Caxton Oluseno, pers. 3300 comm.; Table 4.1). I hypothesise that these may well represent periods of peak 3301 spawning in Lake Chala, although assessment of ovary conditions throughout 3302 the year would be required to corroborate this suggestion. The synchronised 3303 timing of disappearance of both O. hunteri and O. cf. koroque then argues 3304 against a separation in spawning time as a mechanism of reproductive isolation. 3305 However, differential preferences for certain spawning locations could also in-3306 duce a premating reproductive barrier between O. hunteri and O. cf. korogwe. 3307 Different depths of spawning, reinforced by male breeding colour differences, 3308 keep for example O. squamipinnis and O. saka from interbreeding in Lake 3300 Malawi, despite similar overall ecologies (Lowe-McConnell, 1987). 3310

7.3.2 Relationship between tooth shape and feeding habits of Lake Chala tilapiines

Gut content analysis of *O. hunteri*, *O.* cf. *korogwe* and *C. rendalli* over the course of 20 months combined with stable-isotope analysis (Chapter 4) revealed that all three species have a mixed herbivorous-detritivorous diet in Lake Chala,

but display significant seasonal shifts in feeding habits in response to an annual 3316 phytoplankton bloom developing towards the end of the long dry season (July – 3317 September). Nevertheless, evidence for interspecific diet differences was found, 3318 most notably in the stable-isotopic data. In Chapter 2 I hypothesised that 3319 subtle differences in oral tooth shape between O. hunteri and O. cf. korogwe 3320 might facilitate fine-scaled resource partitioning of certain food sources at the 3321 shore (cf. below). However, gut content composition suggested interspecific 3322 divergence in feeding habits to occur mainly during the annual phytoplankton 3323 bloom, when O. hunteri appeared to venture out into the limit zone for 3324 three full months whereas O. cf. koroque (and potentially also C. rendalli) 3325 lingers close to shore until September (Chapter 4). 3326

This seeming mismatch between different oral tooth shapes yet similar 3327 feeding habits during the least favourable times of the year is notable, given 3328 that tooth shape often reflects exploitation of less accessible food sources when 3329 preferred food is unavailable (so-called 'fallback foods'). Perhaps oral tooth 3330 shape of the recently introduced O. cf. korogwe and C. rendalli still reflects 3331 the feeding habits of their respective source populations. Indeed, C. rendalli is 3332 generally a voracious feeder on aquatic and semi-aquatic macrophytes (Ruwet, 3333 1963; Fryer & Iles, 1972), and they are often anthropogenically introduced as a 3334 means of weed control (Dadzie et al., 1988). Considering the almost complete 3335 absence of submerged macrophytes in Lake Chala, it is quite remarkable that 3336 this species has been able to survive and reproduce in the lake for almost half 3337 a century, although the trophic versatility of C. rendalli is well-known. Very 3338 little is known about the feeding habits of O. koroque within its natural range 3339 in the lower Pangani basin (Trewavas, 1983). Insight in the diet of the source 3340 population of O. cf. koroque is moreover complicated by the uncertain identity 3341 of this species in Lake Chala (Chapter 3). The strong (asymmetry linked to) 3342 dental wear which I observed in oral teeth of both O. cf. koroque (Chapter 3343 2) and C. rendalli (data not shown) suggests that their exploitation of Lake 3344 Chala resources comes at a cost. 3345

In this respect, only the oral dentition of *Oreochromis hunteri* probably truly reflects its feeding habits for most of the year, when feeding conditions in Lake Chala are rather harsh in comparison to the short-lived annual phytoplankton bloom. In Chapter 2, I suggested that the slender, shovel-like appearance of oral *O. hunteri* teeth may allow them to function as a comb

and a sieve, ideal for scooping up loose material from between filamentous 3351 epilithic algae covering the rocks ('Aufwuchs', Fryer, 1959a). I did not find 3352 high abundances of macroscopic filamentous algae in the analysed littoral rock 3353 samples (Chapter 4), arguing against frequent combing as a mechanism of 3354 obtaining loose Aufwuchs. Closely-spaced, long and moveable teeth, as found 3355 in small O. hunteri specimens, have also been suggested to enable the removal 3356 of organic detritus from hard substrates in some coral reef fishes (Bellwood 3357 et al., 2014; Hundt & Simons, 2018). Likewise, the detritus, green algae and 3358 diatoms found abundantly in the guts of O. hunteri outside of the annual 3359 period of phytoplankton bloom may have been scooped up from the rocks by 3360 comb-like oral teeth. Nevertheless, the gradual tooth shape shift of O. hunteri 3361 associated with growth (Chapter 6) suggests that there is at least some degree 3362 of change in feeding behaviour linked to ageing (see Section 7.3.3). 3363

The annual phytoplankton bloom in Lake Chala provides highly favourable 3364 feeding conditions for a short period of time (July – September), and the 3365 ability of fish to feed off this bloom probably does not rely on the shape 3366 of their oral teeth. In fact, the capacity of suspension-feeding has in some 3367 Oreochromis species been attributed to a combination of mucus entrapment 3368 and pharyngeal jaw motion, rather than mechanical sieving by gill rakers 3369 (Sanderson et al., 1996). Greenwood (1953) found that a raking motion of the 3370 upper pharyngeal element over the lower pharyngeal element may explain why 3371 large phytoplankton elements occur broken in the gut of O. esculentus. This 3372 is reminiscent of my observations in Chapter 4, where I suggested that the 3373 apparent absence of the large chlorophyte *Treubaria* in fish stomachs, despite its 3374 common occurrence in the phytoplankton, could be due to it being shattened 3375 beyond recognition by the pharyngeal jaws. The interspecific divergence 3376 in fish feeding habits observed during the annual bloom is then potentially 3377 linked to differences in pharyngeal rather than oral tooth shape. An in-depth 3378 morphological study of pharyngeal teeth is outside the scope of this thesis, but 3379 qualitative inspection reveals that pharyngeal teeth of C. rendalli are shaped 3380 very different, and that pharyngeal teeth of O. cf. koroqwe resemble, but are 3381 more robust than, those of O. hunteri (Fig. 7.1). I hypothesise that these 3382 pharyngeal tooth shape differences may compromise consumption of certain 3383 Lake Chala phytoplankton taxa by the introduced tilapiines, and may explain 3384 why O. cf. korogwe and possibly also C. rendalli feed in the littoral area 3385

³³⁸⁶ longer than *O. hunteri*, where their oral dentition guarantees at least some ³³⁸⁷ food intake.



Figure 7.1: Pharyngeal tooth extracted from the lower pharyngeal element of *O. hunteri* (a), *O.* cf. *korogwe* (b) and *C. rendalli* (c).

3388 7.3.3 Ontogenetic shifts in the oral dentition of *O. hunteri*

When developing a modern-day morphological framework of the oral dentition 3389 of O. hunteri (Chapter 6), I found a replacement of bicuspid and tricuspid 3390 teeth by unicuspid teeth in very large individuals, comparable to observations 3391 in other *Oreochromis* species (Trewavas, 1983). This replacement happened 3392 gradually, with both bicuspids and tricuspids displaying a progressively more 3393 dominant major cusp as specimens grow larger, until nearly all teeth are 3394 unicuspid. This gradual ontogenetic change in oral tooth shape may well 3395 reflect changes in microhabitat preference as O. hunteri grow older. Limits on 3396 available sample sizes did not allow us to assess differences in the gut content of 3397 O. hunteri of different body sizes that could possibly be related to a changing 3398 microhabitat preference (Chapter 4), yet such changes have been documented 3399 in other Oreochromis species (Njiru et al., 2004). Fryer & Iles (1972) moreover 3400 indicate that *Oreochromis* species move progressively into deeper water as 3401 they grow. By doing so in Lake Chala, they may encounter altered habitat 3402 conditions that require changes in the shape of the oral dentition for optimal 3403 exploitation. For example, the areal extent of sandy patches between the rocks 3404 may increase at depths with less wave action, and collecting detritus from this 3405 sand likely exerts different forces on oral teeth than scooping up detritus from 3406 rocks surfaces. 3407

The final replacement with unicuspid teeth in the largest adults, however, probably no longer reflects a change in feeding habits. In many cichlid taxa, an entirely unicuspid oral dentition is often associated with piscivory or insectivory

(Fryer & Iles, 1972; Witte & Van Oijen, 1990; Streelman et al., 2003; Burress, 3411 2015) but optimising such feeding habits would require modification of the 3412 entire jaw apparatus, as piscivorous and insectivorous species typically possess 3413 long snouts and gracile jaws (Streelman et al., 2003). Oreochromis species 3414 are rarely entirely piscivorous or insectivorous (Trewavas, 1983), although 3415 juveniles have been observed to opportunistically feed on insects and smaller 3416 fish (De Moor et al., 1986; Njiru et al., 2004). Moreover, none of the O. hunteri 3417 specimens with a unicuspid dentition in my study material contained fragments 3418 of fish, insects or even zooplankton in their gut contents, nor do their isotopic 3419 compositions show evidence of an omnivorous diet. This is despite the presence 3420 of a diverse zoobenthos community inhabiting the mentioned sand patches, as 3421 indicated by fossils preserved in a sediment core collected near-shore (Chapter 3422 5). Thus, the unicuspid dentition of large O. hunteri is not linked to feeding at 3423 a higher trophic level. More likely, the unicuspid teeth are useful in agonistic 3424 behaviour and territorial defence. Oreochromis males usually build bowers 3425 where they display for females and subsequently reproduce (Turner, 1986; 3426 Turner et al., 1991), and a unicuspid dentition enhances the ability to defend 3427 such territories. Indeed, the shift to unicuspid teeth has been documented to 3428 occur in very large, sexually mature *Oreochromis* individuals of several other 3429 species (e.g., O. aureus, O. mossambicus, O. spilurus), especially in males 3430 (Trewavas, 1983). In Lake Chala, this shift probably comes at a considerable 3431 cost in feeding efficiency, as truly unicuspid teeth are far from ideal for feeding 3432 at the rocks or in the sand. However, they may increase breeding success by 3433 enhancing dominance over other males. 3434

³⁴³⁵ 7.4 Taphonomy on fossil cichlid teeth in Lake Chala

³⁴³⁶ 7.4.1 Burial and transport to a core site in the lake's depositional centre

The diet composition of *O. hunteri* and the two introduced tilapiine species indicated that the three larger cichlids in Lake Chala spend most of their time at the lake periphery (Chapter 4). The majority of cichlid teeth thus probably end up in the sediment in close proximity of the fishes' principal habitat at the lakeshore. These fossils mostly represent teeth shed during the process of tooth replacement (approximately every 100 days for tilapiines; Tuisku & Hildebrand,

1994), but may also be the result of the occasional burial of an entire fish carcass 3444 (Elder & Smith, 1988). Nevertheless, cores collected at the depositional centre 3445 also contain abundant cichlid fossils (Fig. 1.5), which may either have been 3446 deposited in situ or were transported to the centre by sedimentation processes. 3447 In order to assess which scenario is most likely, the distribution of cichlid 3448 teeth (both oral and pharyngeal) in sediments deposited near Lake Chala's 3449 periphery is discussed in Chapter 5. I found a strong correlation between the 3450 depth distribution of cichlid teeth or bones and sediment texture: large-grained 3451 (coarse) sediments contain very high numbers of fossil fish remains. The remains 3452 of various groups of zoobenthos, such as ostracods (seed shrimp) and gastropod 3453 molluscs (snails and limpets) were in the same way correlated with larger grain 3454 sizes. I therefore suggested that the high specific density of all these fossils 3455 limits long-distance transport over the flat offshore lake floor. Fish fossils 3456 are thus buried near the lake periphery in relatively coarse sediments from 3457 which the fine and light materials have gradually been removed by resuspension 3458 and winnowing. Likely only an exceedingly small part of the fish fossils is 3459 transported to the depositional centre by repeated winnowing, except that 3460 occasional landslides from a failing underwater slope may transport higher 3461 abundances of fossils to the centre in turbidite flows (Van Daele et al., 2017). I 3462 therefore concluded that most fossil teeth in the offshore sediment record have 3463 been deposited *in situ*. Moreover, although corrosion linearly altered the size 3464 distribution of fossil teeth (Chapter 6), there was no link between median tooth 3465 size and lake level, indicating limited changes in taphonomic effects linked to 3466 lake-level fluctuations. I therefore assume that fossil oral teeth from the six 3467 fossil assemblages studied in Chapter 6 represent the oral dentition of the Lake 3468 Chala fish population at that time, and that they represent it well. 3469

Also, the characteristic basin morphometry of Lake Chala with near-3470 vertical crater walls surrounding a flat profundal lake bottom (Fig. 1.4) ensures 3471 that the distance between the principal location of fish occurrence (littoral) and 3472 the location of final burial in the long fossil record (offshore depositional centre) 3473 has remained relatively stable over time. Only extremely severe lake-level drops 3474 (> 55 m) could have exposed soft sediments at the lake periphery and forced 3475 O. hunteri to move away from the rocky crater walls, closer to the depositional 3476 centre. No indication of such an event was found in the sediment record of the 3477 last 25,000 years, yet it may have occurred briefly during the penultimate glacial 3478

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maximum (\sim 140-128 kyr BP; Moernaut *et al.*, 2010). Lake-level fluctuations 3479 therefore probably did not significantly influence fossil tooth distributions 3480 throughout the sediment record of Lake Chala until at least 160,000 years ago, 3481 when the central ash cones became exposed and provided a second near-shore 3482 environment (Moernaut et al., 2010). This is in stark contrast to what is 3483 expected in more gently-sloping lake basins, where past lake-level fluctuations 3484 not only greatly affect the distance between the locations of fossil deposition 3485 and final burial when past shorelines were displaced (e.g., Danley et al., 2012), 3486 but also local sedimentation dynamics at the burial site. 3487

³⁴⁸⁸ 7.4.2 Tooth preservation in the offshore sediment record

A considerable proportion of fossil teeth retrieved from the 25,000-yr sediment 3489 record of Lake Chala had broken necks (Chapter 2). If most teeth are indeed 3490 deposited in situ and the influence of tooth abrasion during transport is thus 3491 neglegible, these broken teeth suggest the occurrence of other taphonomic 3492 influences after teeth have been buried. The dentin part of fossil teeth becomes 3493 more brittle with age, whereas the enameloid-covered tooth crown is more 3494 resistant, but may eventually also break as its dentin base is being corroded 3495 (pers. obs.). It is known that the mineralisation process differs between 3496 enameloid and dentin (Sasagawa, 1997), which may induce differences in their 3497 preservation potential. The near-perfect state of preservation of $\sim 170,000$ 3498 year-old fossil teeth retrieved from the DeepCHALLA core sequence (discussed 3499 in section 7.5.1), as compared to the near-complete absence of intact tooth 3500 necks older than 8000 yr BP and progressively larger teeth with increasing age 3501 in the last 25,000 years (Chapter 6), indicates that the state of preservation 3502 is not necessarily correlated with the age of the assemblages. These very 3503 old intact teeth ($\sim 170,000$ year BP) were collected from the coarse-grained 3504 base of a turbidite interval, which in Lake Chala have a higher proportion of 3505 clastic material (Van Daele et al., 2017). The relatively lower proportion of 3506 organic material present in these layers may result in lower concentrations of 3507 the organic acids that are produced during decomposition of organic material, 3508 and thus create better conditions for the preservation of fossil teeth, as more 3509 acid conditions could enhance corrosion of calcified fossils such as teeth. The 3510 lack of fossils in one interval studied in Chapter 2 ($\sim 19,200$ yr BP) could thus 3511 also be the result of increased corrosion in this interval. Moreover, the rapid 3512

³⁵¹³ burial of fossil remains at the base of a turbidite may further enhance fossil ³⁵¹⁴ preservation.

³⁵¹⁵ 7.5 Long-term history of *Oreochromis hunteri* in ³⁵¹⁶ Lake Chala

3517 7.5.1 Origin of *O. hunteri* in Lake Chala

Phylogenetic analysis of the mitochondrial DNA (mtDNA) of all Oreochromis 3518 species currently inhabiting the Upper Pangani River system enabled us to 3519 assign O. jipe as the closest living relative of O. hunteri in Lake Chala (Chap-3520 ter 3). As the isolated nature of this lake does not allow direct colonisation 3521 through hydrographic connections, introduction of the first fishes by air is the 3522 only plausible explanation for the colonisation of Lake Chala by O. hunteri. 3523 Transportation of live fishes by birds of prey (potentially carrying mouthbrood-3524 ing females), or natural events such as storms causing 'fish rains' have been 3525 suggested as means of aerial introduction (Elmer et al., 2012), yet very little 3526 empirical evidence for either of these mechanisms exists (Hirsch et al., 2018). In 3527 Chapter 3, I suggested that a riverine ancestral population of O. jipe may have 3528 seeded both Lake Jipe and Lake Chala, and more recently also the reservoir 3529 Nyumba ya Mungu (constructed in 1965; Denny, 1978). This riverine seeding 3530 population likely occurred close to Lake Chala, as smaller distances increase 3531 the probability of a transfer by air. 3532

Notwithstanding its elusive mode of arrival, the ancestral O. hunteri 3533 colonised Lake Chala at least 25,000 years ago, based on its continuous fos-3534 sil record (Chapter 6). Referring to observations in neotropical cichlids in 3535 Nicaraguan crater lakes (Elmer et al., 2010b; Elmer et al., 2012), O. hunteri 3536 could either have arrived soon after the lake's formation approximately 250,000 3537 years ago (Moernaut et al., 2010; Verschuren et al., 2017), as happened in Lake 3538 Xiloá; or Lake Chala may have been uninhabited for long periods of time or 3539 have been inhabited by earlier fish populations that later went extinct, as has 3540 for example been suggested for Lake Apoyeque. I hypothesised in Chapter 3 3541 that the ancestral O. hunteri population, provided with ecological opportunity, 3542 expanded quickly after colonisation, and that the genetic signature of this 3543 expansion could provide us with an estimated time of arrival, as has been 3544

done on cichlids in other lakes (Barluenga *et al.*, 2006; Elmer *et al.*, 2012;
Genner & Turner, 2014). I did not find an unambiguous sign of past population
expansion, but the structure of the haplotype network suggested that the *O. hunteri* population of Lake Chala is relatively ancient.

Discovery of abundant fish fossils in a turbidite ('landslide') layer at a 3549 sub-bottom depth of ~ 133 m during the recent DeepCHALLA drilling appears 3550 to corroborate an ancient origin of the O. hunteri population in Lake Chala. 3551 Based on an age model for the Lake Chala sedimentary sequence proposed 3552 by Moernaut et al. (2010), the landslide occurred at least 170,000 years 3553 ago, which is thus also the minimum age of the fish fossils it dislodged. The 3554 discovery of well-preserved fish fossils at such great depths raises hopes for the 3555 opportunity to study oral tooth shape change over very long time scales. The 3556 DeepCHALLA drilling project eventually acquired a sediment core representing 3557 $\sim 260,000$ years, thus covering most of the lake's history. Although the fossil 3558 cichlid record may be interrupted in some sediment intervals (Chapter 2, and 3559 see also Section 7.4.2), geometric morphometrics would quickly reveal whether 3560 teeth belonged to O. hunteri, and thus provide a very good estimate of the 3561 time of arrival of this species in Lake Chala. Moreover the lake has witnessed 3562 several more (severe) lowstands during this period (Moernaut et al., 2010), and 3563 additional samples of alternating high and low lake-level stands may expand 3564 the observations of changes in tooth type abundance linked to lake-level change 3565 (Chapter 6) on even longer time scales. 3566

³⁵⁶⁷ 7.5.2 Long-term trends in the oral tooth shape of *O. hunteri*

Overall the variation in the shape of oral teeth is highly similar between 3568 all analysed fossil assemblages and the extant fauna (Chapter 2, Chapter 3569 6). Therefore, it is most parsimonious to consider O. hunteri as the only 3570 species to have inhabited Lake Chala throughout the last 25,000 years, at 3571 least before recent anthropogenic introduction of O. cf. koroqwe, C. rendalli 3572 and Haplochromis sp. "Chala". Moreover, the close resemblance of fossil and 3573 modern-day oral teeth indicates that O. hunteri unlikely experienced major diet 3574 shifts, and that besides periods of phytoplankton blooms it has probably always 3575 been feeding in the littoral zone (Chapter 4), with only subtle ontogenetic 3576 changes in feeding behaviour (Section 7.3.3). Qualitative inspection of fossil 3577 pharyngeal teeth (pers. obs.) indicated that these teeth have similarly retained 3578

their slender, unicuspid morphology throughout Lake Chala's fossil record. If 3579 the pharyngeal teeth indeed assist in feeding off phytoplankton (as proposed in 3580 Section 7.3.2), seasonal migration to the limit zone in response to an annual 3581 phytoplankton bloom may have been occurring for at least the last 25,000 years, 3582 similar to what is observed in the lake today. The magnitude of the seasonal 3583 phytoplankton bloom is strongly related to interannual climate variability, 3584 and tends to be less developed during El Niño years (Wolff *et al.*, 2011). 3585 Nevertheless, certain blooming taxa, such as the endemic diatom Afrocymbella 3586 barkeri (Cocquyt & Ryken, 2016), have continuously been present in Lake 3587 Chala during the last 25,000 years (Barker et al., 2011), indicating that this 3588 temporary food source has at least to some extent always been seasonally 3589 available to O. hunteri. 3590

Importantly, although all recovered fossil oral teeth can be assigned to O. 3591 hunteri, I documented recurrent shifts in the proportions of oral tooth types 3592 linked to past lake-level fluctuations (Chapter 6). Apart from an approximately 3593 10% increase in the proportion of unicuspid teeth during low lake-level stands, 3594 there is an overall higher abundance of teeth occupying the section of mor-3595 phospace that is currently taken up by larger fish. However, the enameloid 3596 width of fossil teeth suggested that the median body sizes of fish assemblages 3597 did not increase during lowstands. I therefore surmised that fishes living during 3598 lowstands developed a predominantly unicuspid dentition at smaller body 3599 sizes (and thus probably at a younger age). Below, I elaborate somewhat on 3600 the possible ecological interpretations of my findings, and discuss potential 3601 mechanisms that may be responsible. 3602

3603 Potential mechanisms underlying long-term trends in oral tooth3604 shape

When the lake level of Lake Chala is high, soft-bottom habitats are restricted to 3605 small patches of sand in between rocks, likely slightly increasing in areal extent 3606 at water depths beyond wave-induced turbulence (Chapter 4). In Chapter 3607 1, I hypothesised that low lake levels in Lake Chala may have brought an 3608 extensive soft-bottom benthic habitat (and associated food sources) within 3609 reach at the lake periphery (Fig. 1.6), and that this might be reflected in oral 3610 tooth shape. In this respect, I propose two alternative ecological explanations 3611 for the apparent shift in tooth shapes during past lowstand episodes. 3612

Earlier in this discussion (Section 7.3.3), I suggested that modern-day 3613 O. hunteri may move into progressively deeper water with age, as has been 3614 observed in other Oreochromis species (Fryer & Iles, 1972; Ribbink, 1990), and 3615 that such migration may explain subtle differences in the shape of the oral 3616 dentition linked to age. During low lake level phases, O. hunteri may thus have 3617 encountered soft-bottom habitats earlier in life, in particular a large expanse of 3618 soft sediment. Optimal exploitation of food sources would then have required 3619 changes in oral dentition earlier in life. Alternatively, a shift in oral tooth shape 3620 during low lake-level stands may be an indication of earlier sexual maturation. 3621 Changes in the onset of sexual maturation in *Oreochromis* species are well-3622 described, mostly in the context of aquaculture (Pullin & Lowe-McConnell, 3623 1982; El-Sayed, 2006), and are strongly linked to environmental influences 3624 (reviewed in Brummett, 1995). Changes in the timing of first sexual maturity 3625 have for example been linked to temperature, food availability and quality, and 3626 light regimes, but also to lake-level fluctuations. Specifically in Lake Sibaya 3627 (South Africa), Oreochromis mossambicus starts reproducing at smaller sizes 3628 during episodes of higher lake level, as young males attempt to exploit the 3629 extension of available habitat by breeding in the marginal, recently inundated 3630 regions of the flooded lake (Noakes & Balon, 1982). Although I observed the 3631 opposite signal in Lake Chala (potential reproduction at smaller sizes during 3632 low lake levels), similar underlying mechanisms may be acting. In Lake Chala, 3633 reduced lake levels could increase suitable nesting habitat (i.e. by extension 3634 of sandy substrate), which could allow subordinate (often smaller) males to 3635 start reproducing, thus inducing sexual maturity at earlier ages. Both cues are 3636 not mutually exclusive, and the observed trends in oral tooth morphology in 3637 response to changing lake levels may well be the result of an interplay between 3638 the two. 3639

Based on observations of changing oral tooth shape in the fossil record 3640 of Lake Chala, I hypothesise that a modified developmental trajectory of the 3641 oral teeth enabled O. hunteri to respond to past environmental change (in 3642 the form of major lake-level fluctuations). During low lake-level stands, I 3643 found an increase in the relative proportion of unicuspid oral teeth that was 3644 decoupled from a larger median body size. The developmental trajectory of 3645 O. hunteri teeth may either be altered by increasing the rate of oral tooth 3646 replacement, or by increasing the morphological differences between subsequent 3647

tooth generations, resulting in 'adult-like' dentition occurring in smaller fisheswhen lake levels are low.

Modification of developmental trajectories is a straightforward way to alter 3650 the phenotype, and may provide a valuable mechanism for organisms to rapidly 3651 respond to environmental cues (West-Eberhard, 2003; Shapiro et al., 2004). 3652 Such a mechanism has been suggested to lie at the basis of the distinction 3653 between substrate spawning and mouthbrooding (Noakes & Balon, 1982), and 3654 more recently it has been invoked in the rapid adaptation of cichlids to novel 3655 photic environments upon colonisation of Nicaraguan crater lakes (Härer et al., 3656 2017). Continuous tooth replacement (polyphyodonty) combined with the 3657 potential to rapidly adjust gene expression regulating tooth morphology in 3658 subsequent developmental stages has been suggested to provide cichlids with 3659 a mechanism to quickly adapt to new ecological circumstances (Streelman 3660 et al., 2003). I propose that a similar mechanism has assisted O. hunteri in its 3661 adaptation to major past lake-level fluctuations in Lake Chala, either as a direct 3662 response to climate-driven change in suitable feeding habitat, or indirectly as 3663 the result of climate-driven changes in suitable breeding grounds through a 3664 shift in the onset of sexual maturation. 3665

Long-term trends in oral tooth morphology in response to lake-level fluctuations

I can currently only hypothesise on the exact ecological and evolutionary 3668 mechanisms underlying trends in oral tooth shape of O. hunteri, but the 3669 recurrent shifts in oral tooth shape directly linked to long-term environmental 3670 pacers such as lake-level change are a remarkable observation. They provide 3671 evidence of the ability of *O. hunteri* to rapidly adapt to a changing environment, 3672 and to systematically do so over very long timescales, highlighting the versatility 3673 that typifies the genus Oreochromis. Moreover, rapid phenotypic responses of 3674 the trophic apparatus to environmental change have been observed in other 3675 cichlid lineages (Witte et al., 2008; Muschick et al., 2011; Huysseune, 1995), 3676 and are increasingly accepted as a promotor of diversification (Schneider & 3677 Meyer, 2017). In this thesis, I provided evidence of such responses, and of their 3678 recurrent nature, on very long timescales ($\sim 25,000$ years), indicating that they 3679 may well have enabled cichlids to cope with past environmental change, and 3680 promoted their subsequent diversification and radiation when presented with 3681

³⁶⁸² ecological opportunity.

3683 7.6 Future prospects

Many ecological aspects of modern-day and fossil *O. hunteri* populations have been elucidated in this thesis, yet some important questions remain unresolved. Certain hypotheses that are presented in this Chapter could however be tested in future research.

Elucidating the exact timing of peak spawning of tilapiines in Lake Chala was considered outside the scope of this thesis, but may to some extent influence the migratory behaviour observed in the lake. Breeding season in tilapia species is generally studied by monthly analysis of gonad development in a sufficiently high number of specimens (Babiker & Ibrahim, 1979; Admassu, 1996). Higher frequencies of breeding fish in certain months then indicate periods of peak spawning.

The function of unicuspid teeth in large specimens of O. hunteri is an 3695 important knowledge gap for the conclusions of this thesis. I hypothesised that 3696 the shift towards a more unicuspid dentition is either the result of changes in 3697 feeding behaviour, or linked to sexual maturation and increased territoriality 3698 (Section 7.5.2). To test the former hypothesis, a first step towards a better 3699 understanding of the ecological function of different tooth types would be to 3700 quantify jaw morphology associated with a certain dentition. Indeed, differences 3701 in jaw shape have been linked to tooth shape and body size in other cichlid 3702 species (Streelman et al., 2007), and could give an indication of the predominant 3703 feeding strategy (Kassam et al., 2004). Moreover, morphological comparison 3704 of O. hunteri to other Oreochromis species associated with a soft-bottom 3705 environment could reveal whether the observed shifts in oral dentition indeed 3706 reflect an adaptation to a soft-bottom feeding mode. Truly benthic Oreochromis 3707 species are rare (Trewavas, 1983), but the Nyasalapia species complex (an 3708 Oreochromis subgenus) inhabiting the ecologically diverse Rungwe crater lakes 3709 in Tanzania (Pillay, 2016) provides an ideal control system. These crater lakes 3710 represent a host of different habitats, ranging from entirely soft-bottom to 3711 rocky crater walls interspersed with sandy patches. Quantifying the trophic 3712 morphology of Nyasalapia species associated with each of these habitats may 3713 provide an eco-morphological framework in which the morphology of O. hunteri 3714

³⁷¹⁵ could be interpreted.

Testing whether an earlier shift to a unicuspid dentition is linked to 3716 earlier sexual maturation is ideally done in a controlled setting. However, 3717 differences in the timing of onset of sexual maturation are widely documented 3718 in Oreochromis species grown in ponds (Pullin & Lowe-McConnell, 1982; El-3719 Sayed, 2006). A straightforward approach could be to quantify oral tooth 3720 shape in a representative subset of *Oreochromis* specimens from ponds with 3721 known differences in the timing of onset of sexual maturation, and test whether 3722 unicuspid teeth occur at smaller body sizes in ponds where earlier sexual 3723 maturation occurs. 3724

3725 7.7 General conclusions

This general discussion integrated the results of five chapters describing the extant and past life history of the endemic cichlid *Oreochromis hunteri* in Lake Chala based on its oral dentition, and elaborated on several of my findings. The main conclusions which can be drawn from this doctoral thesis research are summarised below.

- I developed a geometric-morphometric approach for quantitative analysis
 of oral tooth crown shape with landmarks and semi-landmarks. This
 method was highly suitable to distinguish oral tooth shape at the species
 level, yet also for the detection of subtle ontogenetic changes in the extant
 O. hunteri, and of significant shifts in oral tooth shape between different
 fossil assemblages preserved in Lake Chala sediments.
- 2. I found no signs of ongoing hybridisation between the indigenous *O. hunteri* and the recently introduced *O.* cf. *korogwe*, neither in wholebody morphology, nor in two mitochondrial loci. I thus confirmed both the genetic and morphological integrity of the modern-day *O. hunteri* population inhabiting Lake Chala, and its suitability as extant reference framework for fossil tooth assemblages preserved in the lake's sediment record.
- 3744 3. Both *O. hunteri* and the recently arrived tilapiines *O.* cf. *korogwe* and 3745 *C. rendalli* feed most of the year in the littoral zone, but they migrate 3746 seasonally to the limnetic zone to exploit the phytoplankton bloom that

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develops at the end of the long dry season. Although oral tooth shape was 3747 clearly distinctive between species, a significant interspecific difference in 3748 feeding habits appears to occur only during the phytoplankton bloom, 3749 when oral dentition is presumably of less importance. I hypothesise that 3750 the pharyngeal rather than oral dentition may play a role in optimal 3751 exploitation of phytoplankton by Lake Chala tilapiines. As yet undocu-3752 mented shape differences in pharyngeal dentition may thus explain why 3753 the two introduced tilapines have not yet completely adapted to the 3754 seasonally variable food availability in Lake Chala. 3755

4. The distribution of cichlid fossils throughout a core collected near-shore 3756 indicated that the hydrostatic properties of fossil teeth are not favourable 3757 for transportation over very long distances, and suggested that such 3758 transportation is rare. Sedimentological evidence is indicative for in situ 3759 deposition of fossil teeth in offshore locations. 3760

5. Oral tooth shape studied in six fossil assemblages from the last 25,000 3761 years indicate that O. hunteri has been the only species inhabiting Lake 3762 Chala during this period. However, I found clear trends in the shape of 3763 oral dentition linked to major lake-level fluctuations driven by long-term 3764 climate variability. During low lake-level stands, a higher proportion of 3765 unicuspid teeth, typical for large specimens in the modern-day population, 3766 is found. However, this change in dentition was not coupled to increases 3767 in median body size, and thus appears to occur earlier in life. I interpret 3768 this as a change in the ontogenetic trajectory of oral tooth development 3769 in response to changes in feeding or breeding habitats, or a combination 3770 of both. 3771

To my knowledge, I have provided the first evidence of long-term recurrent 3772 changes in the oral dentition of a cichlid species in relation to climate-driven 3773 environmental fluctuations using the fossil record. More generally, I conclude 3774 that building the required control to accurately study individual lineages is 3775 difficult, but the specific setting and well-resolved environmental history of 3776 Lake Chala made it possible for O. hunteri. I anticipate that it will be more 3777 difficult in larger, less isolated lakes which harbour many cichlid lineages, but 3778 a different approach based on the overall tooth diversity at the community 3779 level may be possible there. Moreover, great numbers of teeth are required to 3780

reliable reconstruct patterns through time, and these numbers would need to be further increased for more complex fish communities. Nevertheless, studying a continuous fossil record in lakes with a more complex history and more diverse species flocks may elucidate patterns of long-term eco-morphological change on a larger scale, and further enhance our understanding of processes that underlie cichlid species diversity.

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Supplementary information for Chapter 2

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Table A.1: Number of oral teeth included in geometric morphometric analyses. OH: *Oreochromis hunteri*; OK: *Oreochromis* cf. *korogwe*; DTL1/PML1: outer series; DTL2/PML2: outer row of inner series; DTLA/PMLA: anterior group of the inner series; DTLM/PMLM: middle group of the inner series; DTLP/PMLP: posterior group of the inner series.

]	left denta	ry			le	ft premax	illa		total
	DTL1	DTL2	DTLA	DTLM	DTLP	PML1	PML2	PMLA	PMLM	PMLP	
OH	16	11	33	13	10	34	36	23	18	14	208
OK	19	17	34	14	15	29	20	36	9	16	209

А

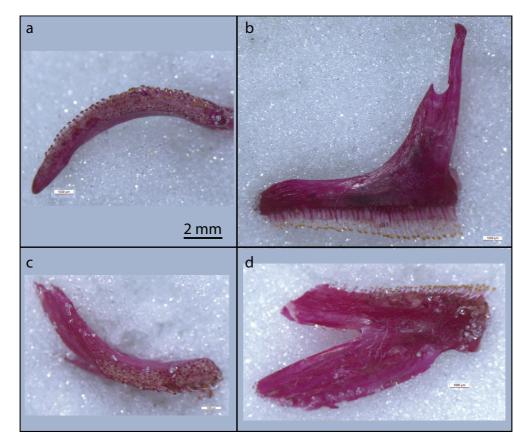


Figure A.1: Overview pictures of *in toto* cleared-and-stained jaws of *Oreochromis hunteri*.a Ventral and b labial view of the right premaxilla, and c dorsal and d labial view of the right dentary.

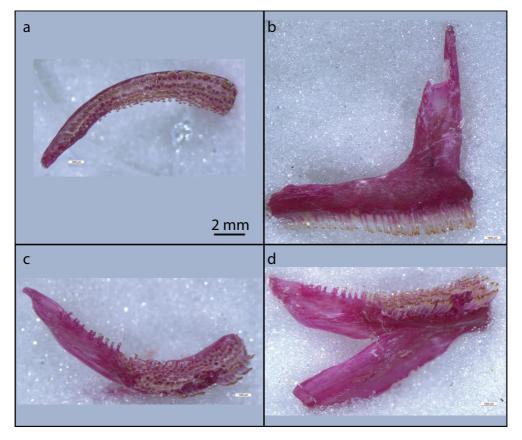


Figure A.2: Overview pictures of *in toto* cleared-and-stained jaws of *Oreochromis* cf. *korogwe.* **a** Ventral and **b** labial view of the right premaxilla, and **c** dorsal and **d** labial view of the right dentary.

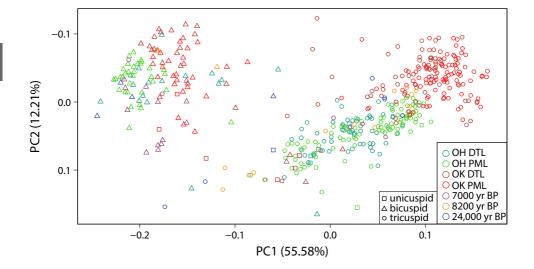


Figure A.3: PCA ordination of SLM-coordinates documenting crown morphology of all modern-day teeth, and fossil teeth of the three age intervals (7,000, 8,200 and 24,000 years ago) analysed. Symbols represent the three different tooth types present in both *Oreochromis* species: unicuspid, bicuspid and tricuspid teeth.

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4656	Supplementary information for Chapter 3
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species	local name	n	n (genotype)	specimen ID	GenBank accession n° CR	GenBank accession n° ND2	locality
Oreochromis hunteri	Chala	81	15	CHALA_cha002	MG922031	MG922069	Lake Chala
				CHALA_cha003	MG922032	MG922070	Lake Chala
				CHALA_cha017	MG922035	MG922073	Lake Chala
				CHALA_cha018	MG922036	MG922074	Lake Chala
				CHALA_cha019	MG922037	MG922075	Lake Chala
				CHALA_cha020	MG922038	MG922076	Lake Chala
				CHALA_cha021	MG922039	MG922077	Lake Chala
				CHALA_cha023	MG922041	MG922079	Lake Chala
				CHALA_cha025	MG922042	MG922080	Lake Chala
				CHALA_cha026	MG922043	MG922081	Lake Chala
Areachramis of karaanne				CHALA_cha027	MG922044	MG922082	Lake Chala
				CHALA_cha028	MG922045	MG922083	Lake Chala
				CHALA_cha029	MG922046	MG922084	Lake Chala
				CHALA_cha030	MG922047	MG922085	Lake Chala
				CHALA_cha043	MG922050	MG922088	Lake Chala
Oreochromis cf. korogwe	Bandia	23	9	BANDIA_cha004	MG922033	MG922071	Lake Chala
Oreochromus cj. korogwe				BANDIA_cha005	MG922034	MG922072	Lake Chala
				BANDIA_cha022	MG922040	MG922078	Lake Chala
				BANDIA_cha031	MG922048	MG922086	Lake Chala
				BANDIA_cha039	MG922049	MG922087	Lake Chala
				BANDIA_cha053	MG922051	MG922089	Lake Chala
				BANDIA.cha054	MG922052	MG922090	Lake Chala
				BANDIA_cha069	MG922053	MG922091	Lake Chala
				BANDIA_cha070	MG922054	MG922092	Lake Chala
Oreochromis jipe	Asilia	18	8	ASILIA_jip003	MG922055	MG922093	Lake Jipe
				ASILIA_jip004	MG922056	MG922094	Lake Jipe
				ASILIA_nym007	MG922063	MG922101	Nyumba ya Mungu
				ASILIA_nym008	MG922064	MG922102	Nyumba ya Mungu
				ASILIA_nym009	MG922065	MG922103	Nyumba ya Mungu
				ASILIA_nym010	MG922066	MG922104	Nyumba ya Mungu
				ASILIA_nym011	MG922067	MG922105	Nyumba ya Mungu
				ASILIA_nym012	MG922068	MG922106	Nyumba ya Mungu
Oreochromis esculentus	Polana	6	6	POLANA_jip005	MG922057	MG922095	Lake Jipe
				POLANA_jip006	MG922058	MG922096	Lake Jipe
				POLANA_nym003	MG922059	MG922097	Nyumba ya Mungu
				POLANA_nym004	MG922060	MG922098	Nyumba ya Mungu
				POLANA_nym005	MG922061	MG922099	Nyumba ya Mungu
				POLANA_nym006	MG922062	MG922100	Nyumba ya Mungu

Table B.1: List of specimens.

species or hybrid	Genbank accession n°
Alcolapia alcalica	AF296455
Alcolapia grahami	KC572512
Alcolapia latilabris	AF296462
Alcolapia ndalalani	AJ237515
Oreochromis amphimelas	AJ237395
Oreochromis amphimelas	AF296490
Oreochromis amphimelas	AF296489
Oreochromis andersonii	JQ907508
Oreochromis andersonii	AF296488
Oreochromis andersonii	AF296487
Oreochromis esculentus	KM654981
Oreochromis esculentus	AJ237399
Oreochromis esculentus	AF296483
Oreochromis esculentus	HQ153030
Oreochromis jipe	AF296492
Oreochromis karongae	AF328844
Oreochromis malagarasi	AF015006
Oreochromis malagarasi	AF296463
Oreochromis malagarasi	AJ237401
Oreochromis mossambicus	AY833447
Oreochromis mossambicus	AF328843
Oreochromis mossambicus	JQ907497
Oreochromis mossambicus	EU430997
Oreochromis mossambicus	AF296466
Oreochromis mossambicus x niloticus	KU180640
Oreochromis niloticus	KU180645
Oreochromis niloticus	HQ153022
Oreochromis niloticus	FJ664216
Oreochromis niloticus	AF296474
Oreochromis niloticus	EU430996
Oreochromis niloticus	AF328847
Oreochromis niloticus	AF296469
Oreochromis niloticus	AF328846
Oreochromis niloticus baringoensis	AJ237397
Oreochromis niloticus vulcani	AF296468
Oreochromis niloticus vulcani	AF296472
Oreochromis niloticus x Oreochromis aureus	FJ664241
Oreochromis sp. HT-1639	AF296486
*	AF296478
Oreochromis sp. HT-1656 Oreochromis sp. HT-8672	AF296470
*	AF296491
Oreochromis sp. HT-9235 Oreochromis sp. HT-9351	AF296464
*	
Oreochromis sp. HT-9352	AF296465
Oreochromis sp. HT-9360	AF296479
Oreochromis sp. KM-2006	AP009126 EU375454
Oreochromis sp. YCC-2008	
Oreochromis sp. YCC-2008	EU375453
Oreochromis spilurus	EU431000
Oreochromis tanganicae	AY929940
Oreochromis urolepis	AF296467
Oreochromis urolepis hornorum	EU308500
Sarotherodon melanotheron	NC015611

Table B.2: Genbank accession numbers for Control Region (CR) sequences.

species or hybrid	Genbank accession n°
Alcolapia alcalica	GQ167781
Alcolapia grahami	KC572546
Oreochromis amphimelas	AF317230
Oreochromis andersonii	AF317231
Oreochromis andersonii	GQ167805
Oreochromis chungruruensis	KF772215
Oreochromis esculentus	KM654981
Oreochromis esculentus	AF317232
Oreochromis karongae	DQ465030
Oreochromis leucostictus	AF317233
Oreochromis leucostictus	KF772225
Oreochromis macrochir	AF317235
Oreochromis mossambicus	DQ465032
Oreochromis mossambicus	AF317234
Oreochromis mweruensis	AF317236
Oreochromis niloticus	AF317237
Oreochromis niloticus	AB018974
Oreochromis niloticus	U07258
Oreochromis niloticus	KF772217
Oreochromis niloticus	GQ167780
Oreochromis niloticus vulcani	AF317242
Oreochromis schwebischi	AF317238
Oreochromis shiranus	KF772216
Oreochromis sp. KM-2006	AP009126
Oreochromis squamipinnis	KF772214
Oreochromis tanganicae	AF317240
Oreochromis tanganicae	KM288941
Oreochromis tanganicae	KM288943
Oreochromis tanganicae	GQ167782
Oreochromis urolepis	AF317239
Oreochromis variabilis	AF317241
Oreochromis variabilis	KM658973
$Sarotherodon\ melanotheron$	NC015611

Table B.3: Genbank accession numbers for ND2 sequences.

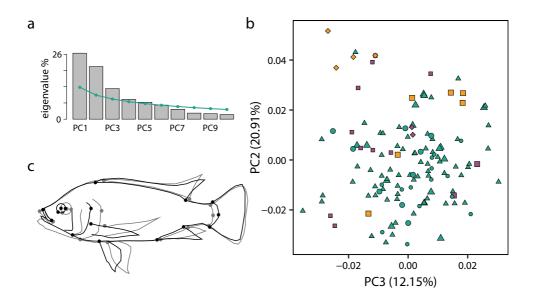


Figure B.1: a Scree plot (grey bars) and broken-stick model (green line) for the first 10 principal components (PCs) capturing body morphology. b PCA ordination plot synthesising variation in body morphology as captured by PC2 and PC3, among the four *Oreochromis* species currently inhabiting the Upper Pangani River basin. Green, purple and yellow symbols represent specimens from respectively Lake Chala, Lake Jipe and NyM reservoir, with large symbols of each type and colour indicating sequenced specimens. c The outline drawings represent the specimens at the positive (black) and negative (grey) extremes of PC3.

С

Supplementary information for Chapter 4

4662 C.1 Supplementary analyses

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We assessed whether gut content composition of each of the three species is influenced by body size. Correlations were tested between NMDS-scores and standard length (SL; distance between the tip of the snout and the posterior end of the last vertebra) for each species separately, using non-parametric Spearman-rank tests. Similarly, we tested for correlations between stableisotopic values (δ^{15} N and δ^{13} C) and SL. None of these correlations yielded a significant result, and we conclude that body size does not influence diet.

	Feb '14 $n = 2$	Mar '14 $n = 3$	Apr '14 $n = 1$	May '14 $n = 2$	u = 114 u = 2	Aug '14 $n = 2$	Sep '14 $\mathrm{n} = 3$	Dec '14 n = 3	Jan'15 n = 2	Feb '15 $n = 2$	Apr $^{7}15$ n = 2	n = 2 Jun ⁷¹⁵	sep '15 $n = 9$
Cyanobacteria	58.1	16.2	47.3	5	4.8	4.3	34.2	3.9	17.7	10	20.7	17.5	73.2
Chroococcales	87.2	24.1		4.2	5.5	4.8	38.1	6.3	34	18.4	31.1	24.3	38.1
Oscillatoriales			64.9							0.8			40.1
Chlorophyta	5.8	6.3	0.5	9.6	52	33.4	32.6	16.1	15.4	31.8	31.3	22.2	17.8
Lagerheimia		0.5		7.3	4.1	0.3		0.3	1.5	2.9	2.1	1.1	1.6
Oocystis	7.6	2.8		1.6	0.9	1.1		0.8	7.5	7.1	2.1	0.4	5.9
Tetraedron	0.6	5.1		9.9	49.6	30.7	34.3	14.9	18	47.7	37.4	27.6	8.7
Treubaria				0.5	0.6	0.3							0.2
other Chlorococcales	0.6	1	0.7	1	4.6	4.2	2	9.9	2.5	3.3	5.5	1.8	1.6
Desmidiales						0.5	0.2	0.3		0.4			1
Dinophyta			1	0.7			0.1	0.5	0.8	0.4	0.6	0.8	0.1
Peridiniales			1.4	1.6			0.2	0.8	1.5	0.8	0.8	1.1	0.1
Bacillariophyta	2.7	44.7	23.2	34.6	30.1	52.5	22.8	41.1	18.2	9.5	13.7	31.5	2.4
A chn anthidium				1									
A frocymbella		0.8	0.7	7.3	0.9	0.8	11.5	1.8	3.5	2.1	2.1	2.2	0.4
Amphora		1.5		2.1		0.3	0.2	1.6		0.4			0.1
Coccone is		0.3										0.4	
Cymatopleura								0.3	0.5	0.4			
Diplone is				1.6	0.6							0.7	
Encyonema													0.1
Encyonopsis				1.6	1.4							1.5	0.3
Navicula		1.3										0.4	
Nitzschia	1.2	44.2	16.9	40.6	18.6	27.5	12.2	45.7	20.5	9.2	9.7	24.3	1.3
Placone is				1									
Rhopalodia		0.8		4.7			0.3	0.3	1	0.4			0.3
Sellaphora				0.5		0.3						1.1	0
Staurosirella	2.3	17.8	14.2	13	0.9	1.3		16.7	9.5	5	7.6	12.9	0.1
Surirella										0.4			
Fragilaria s.l.	0.6				12.2	28	1.2	0.3		0.4	1.3	0.4	0.1
Euglenophyta			1	0.2	0.3			0.2			0.3		0.2
Euglenales			1.4	0.5	0.3			0.3			0.4		0.2
detritus	25.2	17.2	23.2	29.2	9.3	3.6	3.4	25.4	22.4	33.3	16.2	14.3	4.1
sand orains	10	1	0	0.00	1								

С

Table C.2: Numerical abundance (%N) for each algal individual in the combined limnetic samples, collected per month. Bold values represent the relative abundance of phyla, whereas normal fonts are used for relative abundances of algal genera and orders. Supplement to Chapter 4.

	Feb '14			May'14	Jul '14	Aug'14	Sep '14			Apr '15	Jun '15	Sep '15
	n = 2	n = 3	n = 3	n = 3	n = 3	n = 2	n = 3	n = 3	n = 3	n = 1	n = 1	n = 1
Cyanobacteria	59.6	14.8	26.3	16.7	10	34.6	6.7	34.1	52	21.4	49.9	12.7
Chroococcales	5.8	5.5	17.8	3.7	1.9	24.4	3.8	2.3	16.6	1.2	11.1	11.9
Nostocales	1.8	1.4	1	0.1	0.3	7.9	0.5	29	29.8	12.9	26.4	0
Oscillatoriales	52	7.8	7.6	12.9	7.8	2.3	2.4	2.8	5.6	7.2	12.4	0.8
Chlorophyta	28.3	66.5	58.6	70	86.1	61.4	71.3	60.1	44.3	63.6	44.9	85.5
Lagerheimia	0.1	0.7	6.3	9.9	2	0.6	0	10.8	2.8	11.1	10.8	12.4
Oocystis	13.2	26.9	16.5	28.7	1	0.4	0.3	15.9	10.3	19.6	6.2	19.8
Tetraedron	13	20.4	4.3	5.8	42.1	42.7	56.5	6.4	2.9	×	6.5	26.6
Treubaria		0.6	1.8	1.8	21.7			0.1	0.2	2	0.8	1.9
other Chlorococcales	1.9	17.3	28.6	19.3	18.8	17.6	14.5	26.7	27.8	21.1	18.3	21.4
Desmidiales		0.6	1	4.6	0.5	0.1		0.2	0.3	2	2.3	3.3
Zygnematales			0.1									
Dinophyta	1.1	12.9	11.1	7.2	0.2	0	0.1	4.4	1.9	12.8	1	0.3
Peridiniales	1.1	12.9	11.1	7.2	0.2		0.1	4.4	1.9	12.8	1	0.3
${f Bacillariophyta}$	10.6	2.1	2.2	0.9	3.3	3.9	21.5	0.2	0	0.2	0.3	0.4
Aulacoseirales		0.3	1	0.1								
Thalasiosirales				0.2								
A chn an thid ium					0.3							
Amphora			0.1		0.1							
Cymbopleura			0.1									
Encyonema		0.1										0.1
A frocymbella		0.1				0.2	8.2					0.3
Navicula	2.2		0.2	0.1		0.3	3.6					
Nitzschia	8.4	1.1	0.6		0.1		7.5			0.2	0.2	
Rhopalodia								0.1				
Staurosirella		0.2		0.3	1.3	1.3	0.5					
Fragilaria s.l.		0.2	0.1	0.1	1.7	2	1.7	0.1			0.1	
Chrysophyta	0.3	1.3	1.3	2.9	0.3	0.1	0.1	1	1.4	2.1	3.8	0.4
${f E}uglenophyta$	0	2.5	0.5	2.4	0.1	0.1	0.2	0.1	0.3			0.8
Englenales		2.5	0.5	2.4	0.1	0.1	0.2	0.1	0.3			80

	$ m rock \ n=12$	$ ext{sand} ext{n} = 2$
Cyanobacteria	12.9	5.8
Chroococcales	19.4	11.8
Nostocales	0.1	-
Oscillatoriales	2.7	
Chlorophyta	20.3	14.9
Lagerheimia	5.1	
Oocystis	6	3.3
Tetraedron	8.3	20.9
Treubaria	0.5	0.5
other Chlorococcales	13.3	4.7
Desmidiales	1.8	0.9
Dinophyta	0.3	0
Peridiniales	0.6	0
Bacillariophyta	24.5	28.4
A chn anthi dium	0.5	0
A frocymbella	12.7	5.2
Amphora	0.3	3.3
Cymatopleura		0.5
Diplone is		0.5
Ency on opsis	1	
Geissleria		1.4
Gomphone ma	0.4	
Navicula	3.9	3.3
Nitzschia	15.3	29.9
Placoneis	0.1	
Rhopalodia	4.7	10
Stauros irella	0.2	0.5
Fragilaria s.l.	3.1	3.3
Euglenophyta	0.1	
Euglenales	0.1	
detritus	32.1	41.5
sand grains	9.7	9.3

Table C.3: Numerical abundance (%N) for each item in the combined littoral samples, collected in September 2015. Bold values represent the relative abundance of phyla but take into account detritus and sand. Normal fonts are used for relative abundances of algal genera and orders, excluding detritus and sand. Supplement to Chapter 4.

Supplementary information for Chapter 6

4674 D.1 Supplementary analyses

Classification of oral teeth – We examined the robustness of the angle-based 4675 classification method by comparing it to a classification using model-based 4676 clustering applying Gaussian mixture models. We used the package Mclust v.5.3 4677 (Fraley et al., 2012) to fit mixture models with different underlying assumptions 4678 as to the variance-covariance of the data, and we evaluated different clustering 4679 solutions based on a Bayesian information criterion (BIC). This method had a 4680 tendency to overfit the data with multiple clusters that represent biologically 4681 unrealistic scenarios, however, if constrained to fewer clusters (i.e. <5) a 4682 three-cluster model was preferred and the best-fitting three-cluster model was 4683 adopted as a grouping variable. 4684

Results of the two tooth-type classification methods are shown in Figure 4685 D.1a-c (based on intercusp angle) and D.1d-f (based on Mclust modelling). 4686 Teeth assigned to each of the three types occupy highly comparable distributions 4687 in NMDS shape space with either method, although type assignments show 4688 slightly greater overlap when based on the intercusp angle (Fig. D.1a & D.1d). 4689 However, the way in which the three tooth types are distributed in the jaw 4690 differs between the two methods (Fig. D.1b & D.1e). Especially many second-4691 row (SR) tricuspid teeth are classified as unicuspid by the Mclust method. We 4692

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described these particular tricuspids in Chapter 2 as possessing a dominant
middle cusp, causing their general shape to resemble a unicuspid tooth despite
having two clear incisions in the crown's outline. Both methods capture a clear
increasing trend of unicuspid-like teeth in larger specimens (Fig. D.1c & D.1f).
However, tooth type classification based on angles reflects biologically relevant
information better, and this method was therefore retained.

As in the modern-day collection, either classification method produces highly comparable distributions in NMDS shape space when fossil teeth are assigned to each of the three main types of oral teeth (Fig. D.2a & D.2c). Moreover, both classification methods capture the increase in the proportion of unicuspid-like teeth during lowstand episodes (Fig. D.2b & D.2d).

 Table D.1: Overview of fossil oral teeth retrieved from six time windows, representing alternating high- and low lake-level stands

Assemblage	Correct composidepth (e	ite	Approximate age range (yrs BP)	Reconstructed lake level	Wet weight (g)	# oral teeth retrieved	# oral teeth in morphometric dataset
	Top	Bottom					
I	66.5	123	500 - 1000	High	645	148	104
II	239	275	2500 - 3000	High	383	199	162
III	600.5	640	7500 - 8000	Low	491	544	463
IV	787	830.5	10,500 - 11,000	High	399	86	79
V	1328.5	1501	17,500 - 19,000	Low	1890	77	63
VI	1861.5	1959	22,000 - 23,500	High	1224	26	15
Total					5002	1080	886

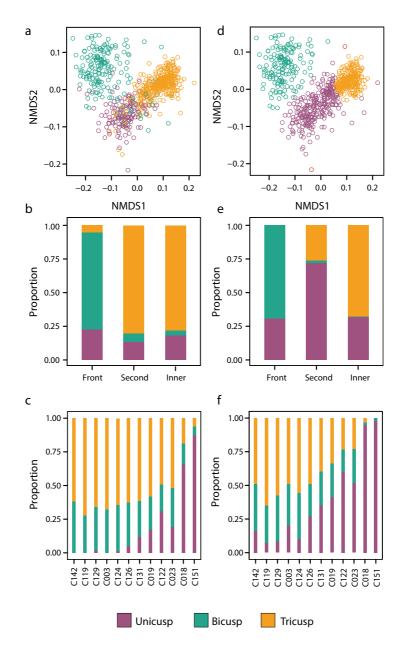


Figure D.1: Results of different tooth-type classification methods in modern-day teeth, with colour codes reflecting unicuspids (purple), bicuspids (green) and tricuspids (yellow). NMDS shape space of tooth types assigned using the intercusp angle (a), and distribution of tooth types in the jaw (b). Proportions of 'traditional' tooth types in each specimen, ordered by increasing body size (c). Tooth type occupation of NMDS shape space based on modelled clustering (d), and their location in the jaw (e). Model-based tooth types per individual (f).

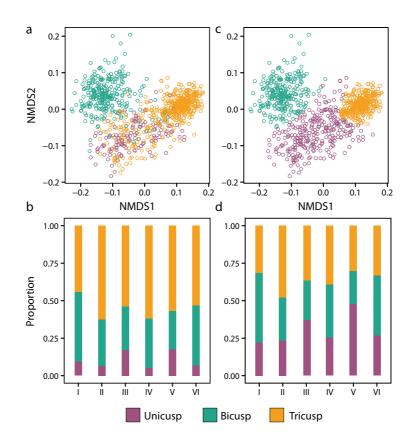


Figure D.2: Results of different tooth type classification methods in fossil teeth, with colour codes reflecting unicuspids (purple), bicuspids (green) and tricuspids (yellow). NMDS shape space of tooth types assigned using the intercusp angle (a), and their relative proportions in the six sampled time intervals (b). Tooth type occupation of NMDS shape space based on modelled clustering (c), and their relative proportions in the six sampled time intervals (b).

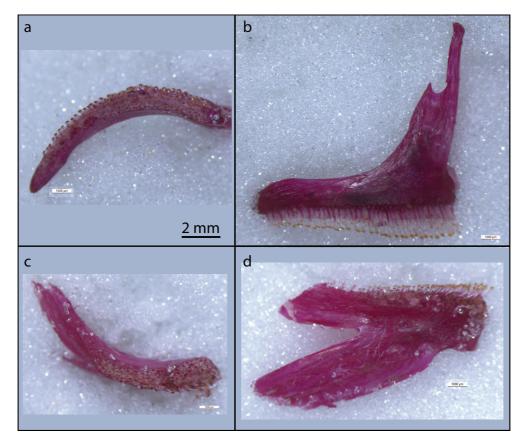


Figure D.3: Overview pictures of *in toto* cleared-and-stained jaws of a small specimen (SL = 18 cm) of *Oreochromis hunteri*. **a** Ventral and **b** labial view of the right premaxilla, and **c** dorsal and **d** labial view of the right dentary.

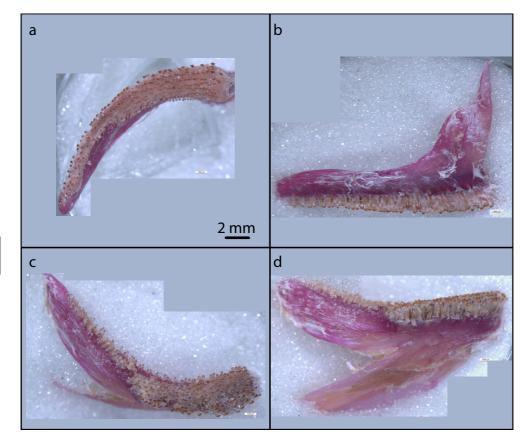


Figure D.4: Overview pictures of *in toto* cleared-and-stained jaws of a medium-sized specimen (SL = 25 cm) of *Oreochromis hunteri*. a Ventral and b labial view of the right premaxilla, and c dorsal and d labial view of the right dentary.

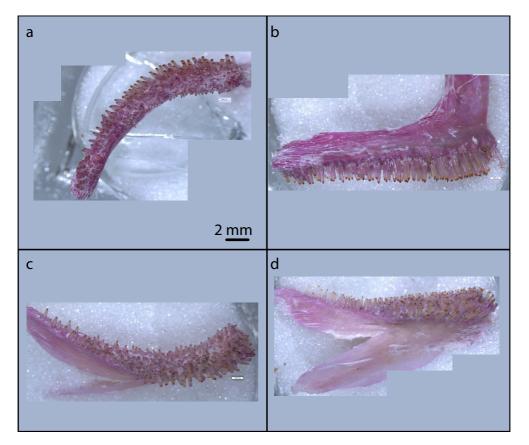


Figure D.5: Overview pictures of *in toto* cleared-and-stained jaws of a large specimen (SL = 30 cm) of *Oreochromis hunteri*. **a** Ventral and **b** labial view of the right premaxilla, and **c** dorsal and **d** labial view of the right dentary. Note the entirely unicuspid dentition of this specimen.