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AN UNUSUAL TOURMALINE COMPOSITION FROM SITHONIA PENINSULA (NORTHERN GREECE)

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Abstract

The variety and wideness of substitutions shown by tourmaline group, mainly in the Z, Y and X sites, make possible a high number of real and hypothetical end-members. However in some cases it is not so obvious to determine to what end-members to refer to, but the wide stability field of tourmaline and its occurrence in different geological environments make useful to define the composition as percentages of end-members to trace back to its genesis. In this paper the followed procedure to identify the components of a complex tourmaline is given. During a study on the granitoid intrusions outcropping on Sithonia, Chalkidiki Peninsula (Northern Greece), a swarm of very small crystals of tourmaline, subhedral or more often allotriomorphic, has been found as stuffing of subparallel micro fractures in the quartz core of a pegmatite dike. Their composition, analysed by EMPA, turns out to be rather unusual and not satisfactorily referable to so far proposed solid solutions. It is characterized by remarkable Al deficiency ($ZAl < 6$ apfu), high amounts of magnesium and iron with $Mg > Fe$ and Na dominating X site. On the basis of compositional and crystal-chemical evidences, supported by previous literature, it was hypothesized that Mg and Fe have to be referred to an uvite – feruvite component, while the considerable part of total iron in its trivalent oxidation state yields a povondraite component according also with Na in X site. These data, coupled with a theoretical evaluation of bond valence sum (BVS) that allowed inferring the presence of OH⁻ anions in the W site, make reasonable to consider Sithonia tourmaline as the new variety sodian povondraite hydroxyl uvite-feruvite. On the basis of its composition, of its fabric and of the oscillatory trends shown by the main elements of Y and X sites, this tourmaline may be considered post magmatic likely of hydrothermal genesis.

Key words: crystal-chemistry, end-members, uvite-feruvite, povondraite, pegmatite.

Περίληψη

Η ποικιλία και το εύρος των υποκαταστάσεων που παρουσιάζονται στην ομάδα του τουρμαλίνης, κυρίως στις θέσεις Z, Y και X, καθιστά δυνατό τον προσδιορισμό ενός σημαντικού αριθμού πραγματικών και υποθετικών ακραίων μελών. Βέβαια σε μερικές

περιπτώσεις αυτό δεν είναι τόσο εύκολο. Ωστόσο, ο προσδιορισμός της % χημικής σύστασής του βοηθά στον καθορισμό παραμέτρων που σχετίζονται με το περιβάλλον γένεσής του. Σε αυτή την εργασία παρατίθεται η μεθοδολογία που ακολουθήθηκε για τον προσδιορισμό των συστατικών μιας σύνθετης μορφής τουρμαλίνη που απαντά στη χερσόνησο της Σιθωνίας. Ο τουρμαλίνης αυτός, που έχει υπιδιόμορφο ή πιο συχνά αλλοτριόμορφο σχήμα, εντοπίστηκε να γειμίζει με πλήθος μικρών κρυστάλλων του παράλληλες μικρο-διακλάσεις στον χαλαζιακό πυρήνα πηγματιτικής φλέβας. Η σύστασή του οδηγεί στο συμπέρασμα ότι πρόκειται μάλλον για μέλος ασυνήθους στερεού διαλύματος. Χαρακτηρίζεται από σημαντικό έλλειμμα Al ($Z_{Al} < 6$ apfu), υψηλές ποσότητες μαγνησίου και σιδήρου με $Mg > Fe$ και Na που υπερτερεί στη θέση X. Με βάση τη σύσταση και το χημισμό του κρυστάλλου, θεωρείται ότι το Mg και ο Fe σχετίζονται με μέλος της σειράς ουβίτη - φερουβίτη, ενώ ένα σημαντικό μέρος του ολικού σιδήρου σε τρισθενή μορφή δίνει ένα ποβονδραϊτικό συστατικό, σύμφωνα επίσης και με το Na στην θέση X. Αυτά τα δεδομένα, σε συνδυασμό με το θεωρητικό υπολογισμό του αθροίσματος των δεσμών σθένους που επιτρέπει τον προσδιορισμό της παρουσίας ανιόντων OH⁻ στη θέση W, μας ενισχύει την άποψη ότι ο τουρμαλίνης της Σιθωνίας ανήκει σε νέα ποικιλία με ονομασία νατρούχος ποβονδραϊτικός υδροζυλ- ουβίτης-φερουβίτης. Με βάση τη σύσταση, τη δομή και τις τάσεις ζώνωσης που παρουσιάζουν τα κύρια στοιχεία των θέσεων Y και X, αυτός ο τουρμαλίνης μπορεί να θεωρηθεί ως μετα-μαγματικής γένεσης, πιθανά με τη συνδρομή υδροθερμικών ρευστών.

Λέξεις κλειδιά: Τουρμαλίνης, κρυσταλλοχημεία, ακραία μέλη, ουβίτης – φερουβίτης, ποβονδραϊτης, πηγματίτης, Σιθωνία, Ελλάδα.

1. Introduction

Tourmaline is a complex cyclosilicate belonging to the hexagonal crystal system, with general formula $XY_3Z_6[BO_3]_3[T_6O_{18}]V_3W$. Its variable chemism, frequent zoning, wide stability field P-T, ability to exchange chemical species with coexisting minerals and fluids (Keller *et al.* 1999) and high strength to exogenous agents make this mineral a potential metallogenic and petrogenetic indicator in many igneous and metamorphic environments (Manning 1982, Henry and Guidotti 1985, Jolliff *et al.* 1986, Kassoli-Fournaraki 1990, Farmer and Halls 1993, London and Manning 1995, Henry and Dutrow 1992, 1996).

Tourmaline structure consists of five types of cationic site: T, tetrahedrally coordinated; B, in triangular coordination; Z and Y, octahedrally coordinated; X, coordinated by nine or ten anions (Selway 1999). Both two octahedral sites and X site are involved in various possible substitutions which give rise to the wide compositional range showed by natural tourmaline.

Usually Al is the dominant cation in Z site, sometimes substituted by even considerable amounts of Mg, Fe²⁺, Fe³⁺ (Fe³⁺₆ = oxy-ferri-foitite and oxy-Mg-ferri-foitite), Cr (Cr₆ = chromdravite and fluor-chromdravite); since substitutions at Y site are constrained more by the size of the ions than by their valence (Grice and Ercit 1993), many different cations may dominate it, that is Li (Li₂Al = liddicoatite and hydroxy-liddicoatite), Mg, Fe²⁺, Fe³⁺, Al and Cr (MgCr₂ = oxy-chromdravite) (Hawthorne and Henry 1999), with lesser amounts of Mn, V, Zn (Schmetzer and Bank 1984, Kazachenko *et al.* 1993).

Several cations may get into X site, mainly occupied by Na, Ca; moreover, it can be K-dominant (Grice *et al.* 1993) or show remarkable vacancies (MacDonald *et al.* 1993, Pezzotta *et al.* 1996, Selway *et al.* 1998a, Aurisicchio *et al.* 1999).

The sites occupied by Si and B are subject to moderate chemical substitutions: small amounts of Al substitute for Si cations in TO₄ tetrahedra (Foit and Rosenberg 1977, 1979, Hawthorne 1996). The presence of B inside TO₄ is debatable, some authors assert there is no evidence for its occupation of Si tetrahedral site, whereas recently Ertl *et al.* (2006) describe ttourmaline

containing IVB from several localities. The presence of substantial IV B is limited to, or more common in, Al-rich tourmalines.

The anions occurring in V and W sites allow defining if the observed phase is an oxy, a hydroxy or a fluoritic term: V site is usually filled by OH⁻, rarely by O²⁻; W site may be occupied by O²⁻, OH⁻ or F⁻.

Two main mechanisms of substitution can occur in tourmaline: the first one involves changes of cations having the same oxidation state inside the same crystallographic site (i.e. ^YMg ↔ ^YFe²⁺); the second one involves changes of heterovalent cations between different sites (i.e. ^XCa + ^ZMg ↔ ^XNa + ^ZAl) (Henry and Guidotti 1985).

Tourmaline composition was currently expressed on the basis of its end-members dravite, schorl and elbaite. However, the analytical method evolution and the new impulse given to researches on this mineral due to its genetic significance showed remarkable deviations and the opportunity to consider new end-members for every possible succession in the three tourmaline groups: alkali, calcic and X-vacant (Hawthorne 1999).

In this work, we discuss the features and the genetic inferences of small tourmaline grains occurring in pegmatitic dykes cutting metamorphic rocks of the Circum-Rhodope Belt, an isopic zone of the internal Hellenides, and related to the nearby granitic intrusions of Sithonia peninsula. The tourmalines studied occur within the dykes as swarms of clusters, composed by small grains or mostly tiny subhedral crystals, showing under microscope evidences of cracks, rarely slight traces of bending, homogeneous distribution of the color (parallel polars) and a sharp pleochroism from pale yellow to deep green. Both the granular aggregates and the subhedral crystals show similar composition.

2. Geological Environment

The Sithonia Peninsula, the middle extension of Peninsula (Northern Greece), is occupied by rocks belonging to the Circum-Rhodope Belt (CRB) and the Serbo-Macedonian Massif (SMM). The CRB is mostly covered by the Sithonia pluton, an elongated Eocene intrusion, and in lesser extent by a Permian to Jurassic flysch of the Svoula Group, by Mid-Jurassic Ophiolites, and by the Chortiatis Unit (Fig. 1) consisting of greenschists, gneisses, metadiorites and marbles (Kauffman *et al.* 1976, Kockel *et al.* 1977). The SMM rocks of the area include the Palaeozoic metamorphic rocks of the Vertiskos Series (Fig. 1) (Kockel *et al.* 1977).

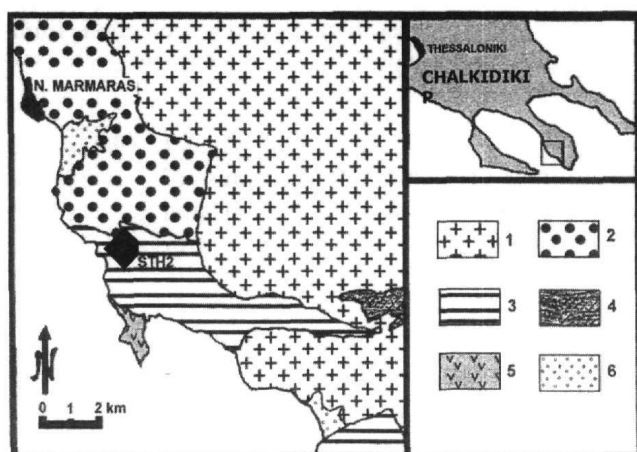


Figure 1 - Schetch map of part of the Sithonia peninsula showing sample (STH2) location. 1. Sithonia pluton; 2. Svoula flysch; 3. Choriatis Unit; 4. Vertiskos Series; 5. Ophiolites; 6. Quaternary.

The Sithonia pluton displays irregular shape and follows the NW-SE regional structural trend. In many places, particularly at the margins of the pluton, a clearly oriented fabric and cataclastic texture can be observed, very probably related to a regional deformative and metamorphic event. The intrusive body comprises moderately peraluminous to metaluminous rocks; the geochemical features suggest processes of fractional crystallization and crustal contamination for the genesis of the Sithonia pluton (Christophides *et al.* 1990).

The ophiolitic rocks and the Chortiatis Unit rocks have been deformed and regionally metamorphosed in the upper greenschist to amphibolite facies mainly during Upper Jurassic to Lower Cretaceous. Retrograde metamorphic events have been also recognized (Mussallam and Jung 1986, Kassoli-Fournaraki 1990).

Aplitic and pegmatitic veins cut both intrusive and country metamorphic rocks.

The samples taken for this study are representative of a dyke cross-cutting gneissic rocks of the Chortiatis Unit in the central area of the peninsula (Fig. 1).

3. Experimental

Chemical compositions were determined in the Sezione di Roma of the Institute of Geosciences and Georesources of the CNR, using a Cameca CX 827 Electron Microprobe, equipped with three wavelength-dispersive (WDS) and one energy-dispersive (EDS) Link-System spectrometers, which are running simultaneously.

Selected crystals have been analyzed for a definite group of elements distinctive of tourmaline, among which only Si, Al, Na, Ca, Mg, Fe, Ti and Mn exceed instrumental detection limits. Compositional homogeneity was verified within single crystals and between crystals belonging to different swarms, to specify eventual chemical variations.

Operating conditions were 15 KV accelerating voltage, 20 nA beam current and 2 μ beam diameter. Wavelength spectrometers were used for all analyzed elements; element peaks and background were measured with counting times of 10 seconds each. A synthetic diopside glass was used as a reference material for Si (TAP), Ca (PET) and Mg (TAP); a jadeite crystal for Al (TAP) and Na (TAP); a rutile crystal for Ti (PET); a Norway fayalite for Fe (LIF); rhodonite for Mn (LIF); a Durango apatite for F (TAP); orthoclase for K (PET). Matrix correction was calculated by using the ϕ (ρZ) method (Pouchou and Pichoir 1984, 1985).

EMP analyses representative of compositions of both granular aggregates and subhedral crystals are reported in Table 1, together with the number of atoms per formula unity (apfu) calculated on the basis of 31 (O, OH), assuming stoichiometric amounts of OH⁻ and B. The assumption B=3 apfu is sustained by several authors (Grice *et al.* 1993, Hawthorne *et al.* 1993, Selway *et al.* 1998) and seems to be satisfactorily correct. Assessment of water content also could be well represented by the formula OH=4apfu (MacDonald *et al.* 1993, Selway *et al.* 1998). The Li content was obtained on the basis of the formula: 3- Σ Y. Fluorine was always under detection limits. Crystallographic sites have been filled partially following previous works (Grice and Ercit 1993, Frietsch *et al.* 1997) and as indicated by some clear correlations.

In order to estimate the Fe²⁺_{tot}/Fe³⁺_{tot} ratio, we imposed the saturation of Si tetrahedral site (Si=6.000 apfu). Using Grice and Ercit (1993) formula the number of Mg atoms at Y site has been calculated; Z site has been saturated to 6.000 apfu with any remaining Mg ions and Fe considered as Fe³⁺ (Frietsch *et al.* 1997).

4. Mineral Chemistry Data and Discussion

The Sithonia tourmaline is characterized by Al deficiency (Al<6.0 apfu), abundance of Mg coupled with high Fe and prevalence of Na in X site; it has an unusual composition, seldom occurring in tourmaline literature, ascribed by some authors to mere dravitic members (Grice and Ercit 1993) or as povondraitic oxy-dravite (Henry *et al.* 1999).

Considering Table 1, Al content is never sufficient to fill the octahedral Z site, whereas Y site occupancy is greater than 3 apfu as was expected considering that some Mg and/or Fe should enter Z site to fill it. Mg ranges from 1.749 to 2.032 apfu, always exceeding Fe²⁺_{tot} which ranges from

Table 1 - Representative EMP analyses of Sithonia tourmaline

	1	2	3	4	5
SiO ₂	36,78	36,45	36,67	36,10	36,49
TiO ₂	0,29	0,00	0,32	0,64	0,22
B ₂ O ₃ *	10,54	10,42	10,48	10,37	10,37
Al ₂ O ₃	29,67	29,32	28,96	28,28	27,70
FeO	7,59	8,76	9,46	10,44	11,37
MnO	0,00	0,15	0,00	0,00	0,00
MgO	8,12	7,51	7,24	7,27	7,45
CaO	0,58	0,61	0,84	0,91	0,54
Na ₂ O	2,61	2,56	2,54	2,44	2,64
Li ₂ O**	0,22	0,19	0,29	0,08	0,07
F	0,00	0,00	0,00	0,00	0,00
H ₂ O***	3,63	3,60	3,61	3,58	3,58
Total	100,04	99,56	100,40	100,11	100,41

Ions number on the basis of 31(O, OH, F)

Si	6,067	6,079	6,084	6,048	6,117
B	3,000	2,999	3,000	3,000	3,000
Al tot	5,770	5,765	5,665	5,586	5,474
Al (T)	0,000	0,000	0,000	0,000	0,000
Al (Z)	5,770	5,765	5,665	5,586	5,474
Al (Y)	0,000	0,000	0,000	0,000	0,000
Fe ²⁺	1,047	1,222	1,312	1,463	1,593
Mn	0,000	0,021	0,000	0,000	0,000
Mgtot	1,996	1,867	1,791	1,814	1,860
Mg(Z)	0,230	0,235	0,335	0,414	0,526
Mg(Y)	1,767	1,632	1,455	1,400	1,334
Ti	0,036	0,000	0,039	0,080	0,027
Li	0,149	0,127	0,192	0,056	0,044
Y tot	2,999	3,001	2,999	3,000	2,999
Z+Y tot	8,999	9,001	8,999	9,000	8,999
Ca	0,103	0,109	0,149	0,164	0,097
Na	0,835	0,828	0,818	0,791	0,856
X tot	0,937	0,937	0,967	0,955	0,953
OH	3,999	4,000	3,999	4,000	4,000
F	0,000	0,000	0,000	0,000	0,000
OH+F	3,999	4,000	3,999	4,000	4,000

Total Iron as FeO.

*calculated on the assumption of 3 B atoms per formula unit.

**calculated as Li = 3-(Ycations sum)

***calculated as OH+F=4

Table 2 - EMP chemical analyses of a single tourmaline crystal

	Rim ext	Rim	Rim	Core	Rim	Rim	Rim opp
SiO ₂	36,50	36,33	36,18	36,35	36,26	36,53	36,27
TiO ₂	0,56	0,52	0,56	0,14	0,12	0,21	0,32
B ₂ O ₃	10,57	10,52	10,48	10,54	10,51	10,58	10,51
Al ₂ O ₃	27,51	27,44	27,50	27,47	27,02	29,71	29,28
FeO*	3,31	3,97	5,07	3,83	1,90	5,32	6,73
Fe ₂ O ₃ *	7,23	6,37	5,34	7,84	9,71	4,00	2,79
MnO	0,00	0,00	0,00	0,00	0,06	0,08	0,11
MgO	7,66	7,74	7,78	7,50	7,26	6,94	7,17
CaO	0,97	0,96	0,96	0,45	0,46	0,60	0,80
Na ₂ O	2,09	2,31	2,31	2,45	2,43	2,40	2,40
Li ₂ O	0,57	0,52	0,39	0,44	0,67	0,43	0,28
F	0,00	0,00	0,00	0,00	0,00	0,00	0,00
H ₂ O	3,65	3,63	3,62	3,63	3,62	3,65	3,62
Total	100,62	100,31	100,19	100,64	100,03	100,47	100,29

Ions number on the basis of 31 (O, OH, F)

Si	6,000	6,001	5,999	5,998	5,998	6,001	6,000
B	3,000	3,000	3,000	3,000	3,000	3,000	3,000
Al tot	5,332	5,343	5,376	5,345	5,269	5,754	5,710
Al (T)	0,000	0,000	0,000	0,000	0,000	0,000	0,000
Al (Z)	5,332	5,343	5,376	5,345	5,269	5,754	5,710
Al (Y)	0,000	0,000	0,000	0,000	0,000	0,000	0,000
Fe ²⁺	0,455	0,548	0,703	0,529	0,263	0,731	0,931
Fe ³⁺	0,894	0,792	0,666	0,974	1,209	0,494	0,347
Mn	0,000	0,000	0,000	0,000	0,009	0,011	0,016
Mg tot	1,877	1,905	1,923	1,845	1,790	1,700	1,768
Mg(Z)	0,668	0,657	0,624	0,655	0,731	0,246	0,290
Mg(Y)	1,208	1,249	1,298	1,190	1,059	1,454	1,478
Ti	0,069	0,065	0,070	0,017	0,015	0,026	0,040
Li	0,374	0,345	0,262	0,290	0,446	0,284	0,188
Y tot	3,001	2,999	3,000	3,000	3,001	3,001	3,000
Z+Y tot	9,001	8,999	9,000	9,000	9,001	9,001	9,000
Ca	0,171	0,170	0,171	0,080	0,082	0,106	0,142
Na	0,666	0,740	0,743	0,784	0,779	0,765	0,770
X tot	0,837	0,910	0,913	0,864	0,861	0,871	0,912
OH	4,000	4,000	4,000	4,000	4,000	4,000	4,000
F	0,000	0,000	0,000	0,000	0,000	0,000	0,000
OH+F	4,000	4,000	4,000	4,000	4,000	4,000	4,000

B₂O₃, Li₂O and H₂O as in table 1.

*FeO and Fe₂O₃ calculated imposing Si=6.

1.050 to 1.597 apfu. Mn is usually under the instrumental detection limits, while Ti sometimes reaches 0.14 apfu. Calculated Li is never greater than 0.15 apfu, as expected for ferro-magnesian compositions.

Tetrahedral T-site always shows a considerable excess of Si cations (Table 1); this repeated result has been confirmed by rigorous checks through the execution of numerous analyses over different crystals and within single crystals and using different reference standards, in order to display Q-micro-inclusions effect and eventual instrumental errors. This excess, as consequence, could be considered consistent with the existence of trivalent cations in the tourmaline structure.

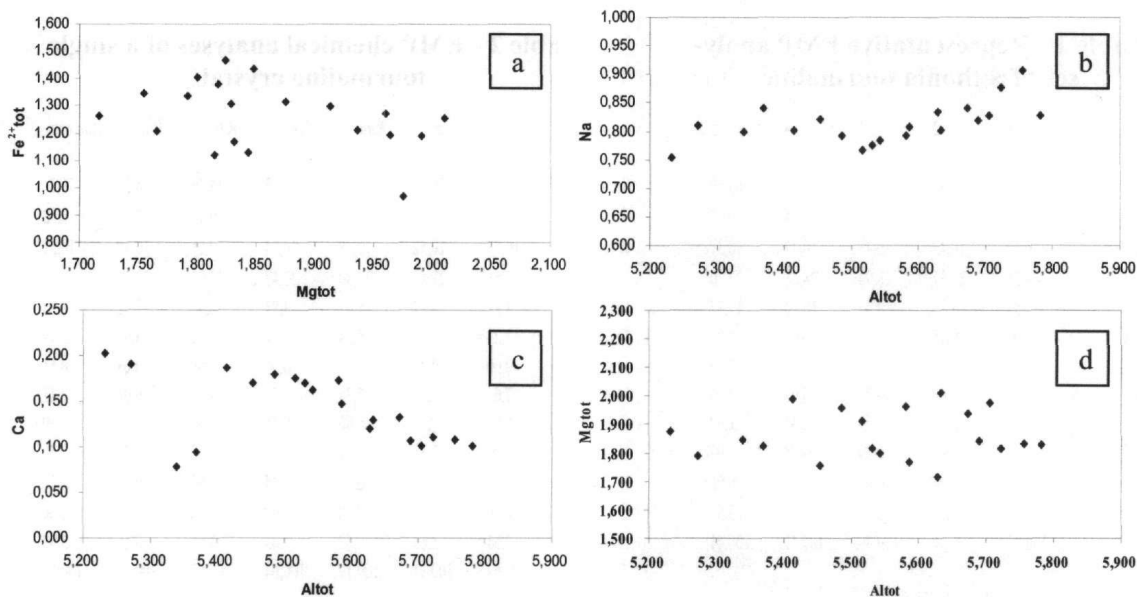


Figure 2 - Linear correlations between the main elements (apfu) of Sithonia tourmaline: (a) Fe^{2+}_{tot} vs. Mg_{tot} ; (b) Na vs. Al_{tot} ; (c) Ca vs. Al_{tot} ; (d) Mg_{tot} vs. Al_{tot}

X-site is dominated by Na, followed by lesser amounts of Ca. Small vacancies are frequent.

The following linear correlations between the main elements allow to visualize the Sithonia tourmaline behavior: Fe^{2+}_{tot} vs. Mg_{tot} shows oscillations around a trend slightly negative (Fig. 2a); a light positive correlation links Na to Al_{tot} (Fig. 2b), whereas Ca and Al_{tot} give a good negative correlation (Fig. 2c); moreover, Mg shows steady values for all Al_{tot} contents (Fig. 2d).

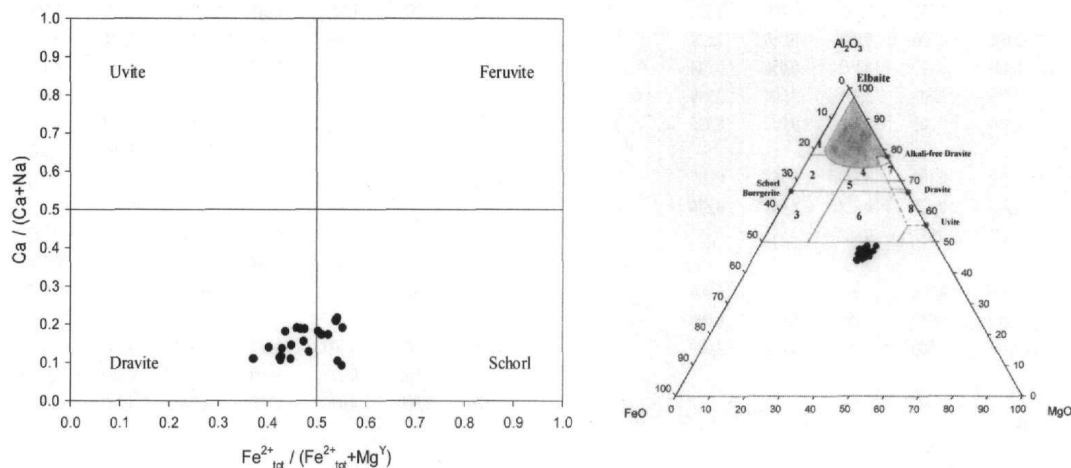


Figure 3 - (a) Plot of $Ca/(Ca+Na)$ at the X site vs. $Fe/(Fe+Mg)$ at the Y site. (b) Al – Fe_{tot} – Mg diagram (in molecular proportions) for tourmaline from various rock types: field #6 defines the compositional range of tourmaline from Fe³⁺-rich quartz-tourmaline rocks, calc-silicate rocks and metapelites (Henry and Guidotti 1985)

The diagrams of Fig. 3 (a and b) represent an attempt for the classification of Sithonia tourmaline on the basis of the correlation $Ca/(Ca+Na)$ vs. $Fe_{tot}/(Fe_{tot}+Mg)$ given by Selway *et al.* (1998) and the compositional fields defined in the triangular plot $Al_2O_3 - MgO - FeO$ proposed by Henry and Guidotti (1985). In Fig. 3a the studied compositions fall on the boundary line between schorl and dravite end-members, with a light prevalence of dravite. The triangular diagram of Fig. 3b shows the area covered by Sithonia compositions: they fall below the $Al_{50} - (Fe, Mg)_{50}$ line.

However, the reported graphic representations of the chemical compositions of Sithonia tourmaline do not seem to be fully representative of these phases, since do not show meaningful trends and do not allow to point out the Al behavior, nor to define any specific field for such compositions. Nevertheless, the triangular diagram of Fig. 3b suggests the presence of a possible considerable Fe^{3+} amount, since they fall just below the field #6 of the triangle reported by Henry and Guidotti (1985) and assigned to Fe^{3+} -rich quartz-tourmaline rocks, to calcium-silicate rocks and to metapelite. This hypothesis may be further confirmed by the optical observations carried out on our samples. In fact, tourmaline containing significant amounts of Fe with mixed valences ($2+$ and $3+$) exhibits very strong optical absorption and pleochroism due to intervalence charge transfer (Mattson and Rossman 1987), as in the case of Sithonia tourmaline.

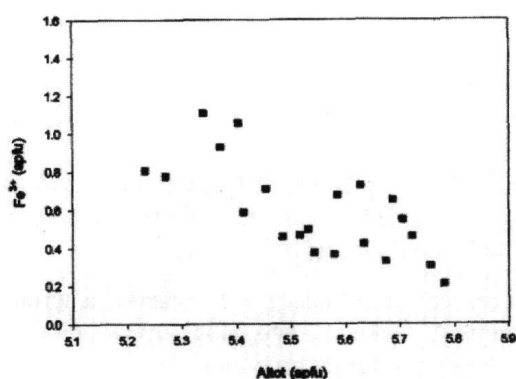


Figure 4 - Tendency of calculated Fe^{3+} vs. Al_{tot} showing a negative correlation (apfu)

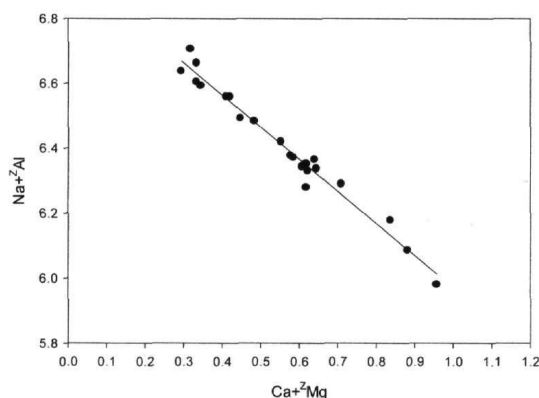


Figure 5 - Negative correlation $\text{Na}+\text{Al}(\text{Z})$ vs. $\text{Ca}+\text{Mg}(\text{Z})$, representing the uvitic substitution (apfu)

To quantify the amount of Fe^{3+} the $\text{Fe}^{2+}_{\text{tot}}/\text{Fe}^{3+}_{\text{tot}}$ ratio was evaluated, using an iterative program, which carries out Fe_{tot} division imposing a stoichiometric amount of $\text{Si}=6.000$ apfu. In Table 2 the analyses of a single tourmaline crystal have been recalculated following the above procedure.

Using the formula proposed by Grice and Ercit (1993), based on the hypothesis of an ideal solid solution at Y site between Fe and Mg for tourmaline compositions with FeO_{tot} greater than 7 %, it was possible to determine the number of Mg atoms at Y site, with any remaining Mg assigned to Z. Often the calculated $^{\text{Z}}\text{Mg}$ is not enough to fill Z site and vacancies continues to exist which have been saturated with the necessary amount of Fe^{3+} to reach 6.000 apfu.

The goodness of this choice, to partially fill Z-site with Fe^{3+} , is confirmed by the good negative trend Fe^{3+} vs. Al_{tot} (Frietsch *et al.* 1997) of Fig. 4; the excellent negative trend of Fig. 5 shows the uvitic substitution $\text{Na}+\text{Al} \leftrightarrow \text{Ca}+\text{Mg}$. Moreover, the Manning (1982) diagram points out transitional compositions toward uvite end-member perfectly following substitution mechanism line (Fig. 6). So, the seemingly simple solid solution schorl-dravite reveal the existence also of an

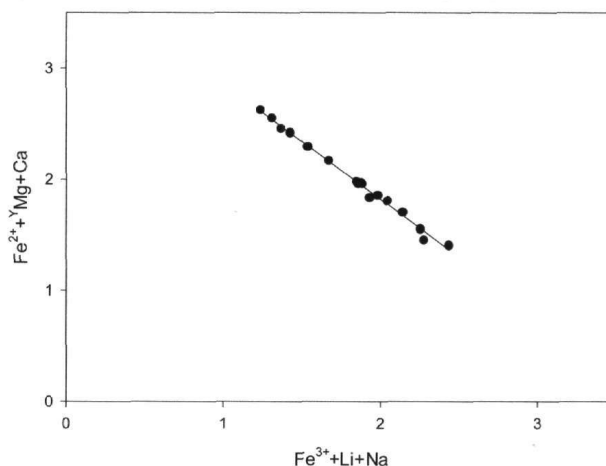


Figure 6 - The complex substitution mechanism $\text{Fe}^{3+}+\text{Na}+\text{Li} \leftrightarrow \text{Mg}(\text{Y})+\text{Ca}+\text{Fe}^{2+}$, controlling the transition povondraite \rightarrow uvite-feruvite (apfu)

uvitic component, as from the entrance of Ca substituting Na in X site proceeding toward lesser Al contents, with always high Mg_{tot} .

Assuming the Al amount as indicator of compositional and thermal evolution, it is known that Al-rich tourmaline crystallize during magmatic and pegmatitic stages, while Al-poor terms have been observed in the later hydrothermal stage (London and Manning 1995, Trumbull and Chaussidon 1999), Sithonia tourmaline crystallized in a Fe-Mg rich and Al-poor environment, according with the fluorine content always under the electron microprobe detection limits, which, coupled with the decrease of T, should expand the stability field of tourmaline toward lower B and Al content (London 1999).

The recorded Fe^{3+} contents are variable but increasing for decreasing Al (Table 2), therefore while tourmaline growth evolves from schorl-dravite toward the uvitic end-member, the oscillatory trend of Fe^{3+} -rich component is the result of a fluctuating increase of the oxydation degree of hydrothermal fluids. Because of the increase of oxygen fugacity, the transition from simple schorl-dravite to more complex solid solution could be combined also with the substitution of OH^- with O^{2-} in W site. A theoretical evaluation of Bond Valence Sum (BVS), utilizing ionic radii values of Shannon (1976) and Brown and Wu formula (1976), with the updating of Brown (1987), allows confirming the predominant OH^- presence in W site, consistently with the results obtained by Grice and Ercit (1993) for similar tourmaline compositions on the basis of X-ray data.

In order to emphasize the significant chemical features of these phases, we consider a ternary diagram functional to support the new reading-key, that is $Al_6 - Fe^{3+}_6 - Mg_6$ (Fig. 7a), which three vertices describe Z occupancy: the first Al_6 indicates the Al saturation amount both in dravite and schorl end-members; the second Fe^{3+}_6 the Fe^{3+} saturation amount in an oxy-Mg-ferriferite and an oxy-ferriferite, while Mg_6 is the Mg amount in a hypothetical term in which Mg entirely fills Z-site. For the sake of clearness, in the same diagram some Z-compositions are plotted, representative of tourmaline end-members particularly useful to display Sithonia tourmaline behavior: the representative points of Sithonia tourmaline fall near the line jointing the top vertex Al_6 with $MgAl_5$ point corresponding with Z occupancy of (hydroxy)uvite, (hydroxy)feruvite and oxy-dravite. The diagram shows that the Al deficiency causes a shift towards "uvite" composition, because the existence of an oxy-dravitic component can be excluded owing to Al(Y) absence. Moreover, the F absence in the analyses leaves out $MgAl_5$ point represents compositions with F anions in W-site (i.e. ideal uvite and feruvite).

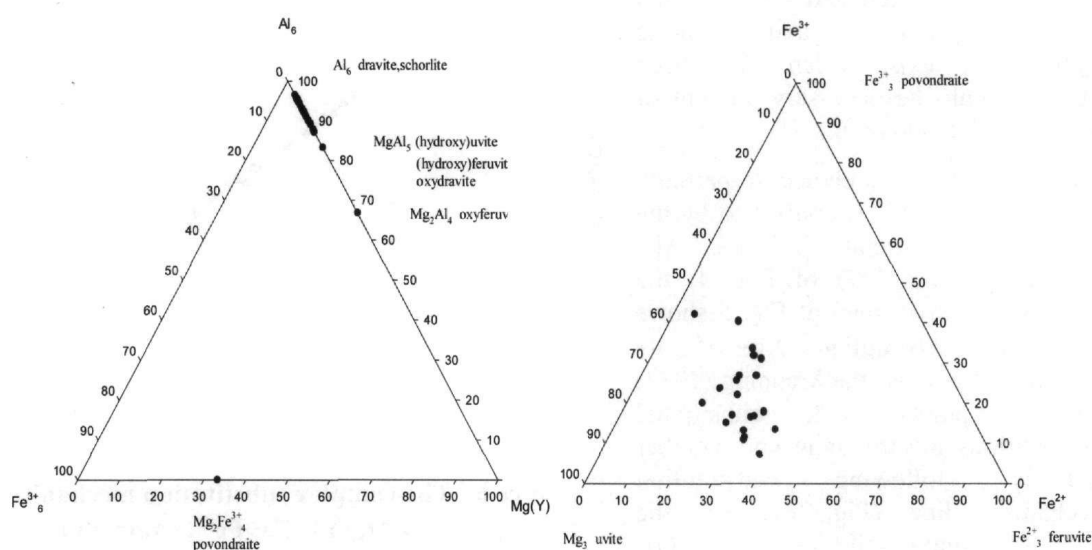


Figure 7 - (a) Triangular diagram $Al_6 - Fe^{3+}_6 - Mg_6$, describing Z-site occupancy, (b) Triangular diagram $Fe^{3+}_3 - Mg(Y)_3 - Fe^{2+}_3$, showing the coexistence of the three components

In the light of the above, we can consider Sithonia compositions as schorl-dravite containing an uvitic and a Fe^{3+} -rich component. The triangular diagram of Fig. 7b, which vertices represent the saturation amounts of Fe^{3+} (Fe^{3+} -rich end-member), Mg (dravite) and Fe^{2+} (schorl) in Y-site, defines the compositional field of this solid solution.

The small dimensions of tourmaline crystals, together with their exclusive occurrence as subparallel veinlets, as well as their complex composition characterized by $\text{Fe}_{\text{tot}}/\text{Mg}_{\text{tot}}$ ratio < 1 (Power 1968, Fieremans and De Paepe 1982, Taylor and Slack 1984, London and Manning 1995, Trumbull and Chaussidon 1999), suggest that tourmaline crystallized during a hydrothermal stage, in subsolidus conditions, as confirmed by $^{\text{Y}}\text{Al}$ absence (London and Manning 1995) and by Al deficiency in Z site. The low degree of alkali-site vacancies (Trumbull and Chaussidon 1999) allows drawing similar inferences on the crystallization of Sithonia tourmaline.

5. Conclusions

- Compositions as Sithonia tourmaline, before considered as simple dravite-schorlite solid solutions or povondraite oxy-dravite, turn out to be a more complex phase, characterized by deficiency of Al in Z-site and predominance of Mg_{tot} over Fe_{tot} , with Na over Ca in X-site. This composition seems to refer to an uncommon combination of several actual and hypothetical end-members.
- The hypothesis of a hydrothermal origin is further on supported by the finely oscillatory trends, distinctive of a growth from rapidly changing fluids. The several peaks of Fe^{3+} within its negative oscillatory trend versus Al_{tot} could be referred to different and more and more oxydized fluid rising (Cavarretta and Puxeddu 1990).
- The variations of the chemical-physical parameters produced a fine-scale zonation of crystals, reflected by continual trend inversions (London 1999). At this state, lacking isotopic studies, it is not possible any claim about the origin of fluids: true metamorphic or with a magmatic component. However, a strong metamorphic contribution seems to be preferred on the basis of micro-structural and geological observations, i.e. the bending traces under the microscope and the gneissic rocks surrounding the quartz-tourmaline dyke; moreover, we cannot avoid to remark that hydrothermal tourmaline, containing high Mg, usually occurs in veins hosted in metasedimentary rocks (Smith and Yardley 1995).

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