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Order of ^[6]Ti⁴⁺ in a Ti-rich calcium amphibole from Kaersut, 39 Greenland: a combined X-ray and neutron diffraction study 40 41 G. Diego Gatta¹, Garry J. McIntyre², Roberta Oberti³, Frank C. Hawthorne⁴ 42 43 ¹Dipartimento di Scienze della Terra, Università degli Studi di Milano, 44 Via Botticelli 23, I-20133 Milano, Italy 45 ²Australian Nuclear Science and Technology Organisation, New Illawarra Road, 46 Lucas Heights NSW 2234, Australia 47 ³CNR-Istituto di Geoscienze e Georisorse, Via Ferrata 1, I-27100 Pavia 48 ⁴Department of Geological Sciences, University of Manitoba, Winnipeg, Canada R3T 2N2 49

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52 Abstract

In order to characterize the role of Ti in the crystal structure of calcium amphiboles with high 53 or even dominant oxo-component, the crystal structure of a Ti-rich calcium amphibole from a gabbro 54 at Kaersut, Greenland, has been refined with single-crystal MoK α X-ray intensity data to an $R_{l}(F)$ 55 index of ~ 0.025, and with single-crystal Laue neutron intensity data to an $R_{I}(F)$ index of ~ 0.053. The 56 crystal used for X-ray structure refinement was characterized by electron- and ion-microprobe analysis. 57 The site populations of the C-group cations Mg, Fe and Ti were calculated from the refined site-58 scattering values for the M(1), M(2) and M(3) sites derived by both X-ray and neutron diffraction. Ti is 59 distributed among all the three 6-fold coordinated M sites, with a strong preference for the M(1) and 60 M(3) sites, where its main role is maintaining electroneutrality at the deprotonated O(3) site. The 61 pattern of distortion of the M(1), M(2) and M(3) octahedra differs from that in F-free deprotonated 62 or partly deprotonated amphiboles, where Ti^{4+} does not occur at the M(3) site. The neutron structure 63 refinement provides also a clear picture of the environment of the proton, anisotropic-displacement 64 behaviour and potential hydrogen-bonding arrangements. A trifurcated hydrogen-bonding 65 configuration has been identified, with two O(6) and one O(7) oxygen atoms as *acceptors* of weak 66 67 hydrogen-bonds.

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Keywords: Ti-rich amphibole, kaersutite, single-crystal X-ray diffraction, single-crystal Laue
 neutron diffraction, crystal chemistry.

Introduction 72

The amphiboles are probably the mineral supergroup with the highest chemical variability 73 74 and the highest ability to crystallize under almost all the conditions relevant to petrogenesis. The general amphibole formula is written (Hawthorne and Oberti 2007) as 75

A₀₋₁ B₂ C₅ T₈ O₂₂ W₂, 76 $A = Na^{+}, K^{+}, \Box, Ca^{2+}, Li^{+};$ where: 77 $B = Na^+, Li^+, Ca^{2+}, Mn^{2+}, Fe^{2+}, Mg^{2+};$ 78 $C = Mg^{2+}, Fe^{2+}, Mn^{2+}, Al^{3+}, Fe^{3+}, Mn^{3+}, Ti^{4+}, Li^+;$ 79 $T = Si^{4+}, Al^{3+}, Ti^{4+};$ 80 $W = (OH)^{-}, F^{-}, Cl^{-}, O^{2-}.$ 81

82 With regard to the present work, C defines the cations occupying the octahedrally coordinated M(1), M(2) and M(3) sites and T defines the cations occupying the tetrahedrally coordinated T(1) and T(2)83 sites. The amphibole structure (with site nomenclature) is shown in Figure 1. Recent studies have 84 shown that the oxo component, ${}^{WO^{2-}}$ or ${}^{O(3)}O^{2-}$ to identify the relvant site, is particularly important in 85 affecting both the geometry and the crystal-chemical behaviour of the amphibole, to the extent that 86 oxo-dominant compositions are now classified as a separate group by the current nomenclature 87 88 scheme (Hawthorne et al. 2012).

The ability of the amphibole structure to incorporate cations in different coordinations and at 89 90 different structural sites, particularly where related to coupled substitutions, makes this mineral supergroup the ideal case study to test crystal-chemical mechanisms controlling the incorporation of 91 Ti⁴⁺ in silicates. 92

93

Previous work 94

Leake (1968) compiled a large number of amphibole analyses from the literature and classified 95 them as to quality. Saxena and Ekström (1970) examined 639 analyses, all first-class and some second-96 97 class analyses as classified by Leake (1968), by principal component analysis. The strongest correlation between chemical variables in this dataset is between Ti⁴⁺ and (OH)⁻. They proposed that 98 Ti⁴⁺ enters calcium and sodium-calcium amphiboles via the substitution 99 **m**•4+ α^2 [6] + 12+ (1) 100

$$11^{4'} + 0^{2'} \rightarrow {}^{[0]}Al^{5'} + (OH)^{2'}$$
 (1)

and found no correlation between the oxo component and the Fe^{3+} content of the amphibole. 101

Kitamura et al. (1975) and Pechar et al. (1989) showed by neutron diffraction that Ti is a C cation and is very strongly ordered at M(1) in two kaersutites. Oberti et al. (1992) showed by X-ray structure refinement that ^[6]Ti⁴⁺ is incorporated into the amphibole richterite structure *via* three distinct substitutions:

 ${}^{M(2)}\text{Ti}^{4+} + 2 {}^{T(1)}\text{Al}^{3+} \rightarrow {}^{M(2)}(\text{Mg},\text{Fe})^{2+} + 2 {}^{T(1)}\text{Si}^{4+}$

$${}^{M(1)}\text{Ti}^{4+} + 2 {}^{O(3)}\text{O}^{2-} \rightarrow {}^{M(1)}\text{Mg}^{2+} + 2 {}^{O(3)}(\text{OH})^{-}$$
 (2)

$$^{T(2)}\mathrm{Ti}^{4+} \to ^{T(2)}\mathrm{Si}^{4+} \tag{4}$$

In richterites from lamproites, the partitioning between the C and T cation-groups depends on the 109 pressure of crystallization (Konzett 1997; Konzett et al. 1997) and the F content of the amphibole 110 (Della Ventura et al. 1991, 1993). In an examination of a series of synthetic Ti-rich Fe-free kaersutites 111 by X-ray diffraction, Tiepolo et al. (1999) showed that ^[6]Ti is disordered over M(1), M(2) and M(3)112 with a strong preference for the M(1) site, and proposed two separate mechanisms for incorporation of 113 ^CTi into the amphibole structure: The first mechanism couples ^[6]Ti to the occurrence of O²⁻ at O(3) and 114 leads to complete order of ^[6]Ti at M(1) (*i.e.*, substitution (2) above); the local association of M(1)Ti⁴⁺ 115 with 2 ^{O(3)}O²⁻ was confirmed spectroscopically by Della Ventura et al. (2007). The second substitution 116 couples ^[6]Ti to the occurrence of Al at the *T* sites (*i.e.*, substitution (5) below): 117

118 ${}^{M(2,3)}\mathrm{Ti}^{4+} + 2 {}^{T(1,2)}\mathrm{Al}^{3+} \rightarrow {}^{M(2,3)}\mathrm{Mg}^{2+} + 2 {}^{T(1,2)}\mathrm{Si}^{4+}$ (5)

augmenting the local association of other small high-valence cations (*e.g.*, Al, Fe³⁺, Cr³⁺) at M(2) and M(3) with Al at T(1) and T(2) in accord with the valence-sum rule (Brown 1981, 2002; Hawthorne 1997; Hawthorne et al. 1998). Della Ventura et al. (1996) showed that in substitution (4), there is no short-range order involving ^[4]Ti and Si.

Therefore, the incorporation and site ordering of Ti in amphiboles is a function of both amphibole composition and conditions of crystallization. Understanding the crystal-chemical constraints on ^CTi partitioning in amphiboles is of great interest as the common occurrence of the oxo component in amphiboles is becoming more apparent due to advances in the structural and analytical techniques and advancing crystal-chemical knowledge. The present work was designed to examine rigorously the ordering of Ti in a calcium amphibole for which comparative crystal-chemistry, based on X-ray diffraction data, indicates that ^CTi is distributed over all the octahedrally coordinated sites.

130

131 Sample provenance

(3)

One sample was selected from the crystal-chemical database of the Pavia unit of the CNR-IGG 132 institution, which contains the results of complete (EMP+SIMS) chemical analysis, X-ray crystal-133 structure refinement and in some cases Mössbauer spectroscopy. This sample was chosen because of 134 its composition (*i.e.*, high Ti content) and the availability of large (centimetric) crystals suitable for 135 neutron diffraction. The sample comes from a gabbro at Kaersut (Greenland), the type locality of 136 kaersutite. The sample code at the Muséum National d'Histoire Naturelle in Paris is MHN109.991, and 137 the crystal code in the CNR-IGG database is 632. Previous work based on comparison with X-ray 138 refinement of synthetic high-Ti Fe-free kaersutite (Tiepolo et al. 1999) had suggested that also this 139 sample had a significant amount of Ti at the M(3) site, and that it cannot be considered as an oxo-140 141 amphibole (namely, kaersutite) because its (OH+F) content, as measured by SIMS, is higher than 1.0 142 atoms per formula unit (apfu).

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- 144

145 **Experimental**

146 - X-ray diffraction

147 Some fragments of a centimetric crystal ($\sim 0.8 \text{ cm}^3$) of the amphibole sample from Kaersut, 148 here labelled as "K(1)", were used for this multi-methodological study.

Unit-cell dimensions of crystal K(1) were calculated from least-squares refinement of the d 149 150 values obtained from 50 rows of the reciprocal lattice by measuring the centroid of gravity of each reflection in the range $-30 < \theta < +30^{\circ}$. Intensity data were collected for the monoclinic equivalent 151 pairs (*hkl* and *h-kl*) in the range $2 < \theta < 65^{\circ}$. Intensities were then corrected for absorption, Lorentz 152 and polarization effects, averaged and reduced to structure factors. Unit-cell dimensions and 153 154 miscellaneous information pertaining to intensity-data collection and refinement are given in Table 1. Only reflections with $I_0 > 3\sigma_I$ in the range $2 < \theta < 30^\circ$ were considered as "observed" during 155 unweighted full-matrix least-squares refinement on F, done with an extensively modified version of 156 the program ORFLS (Busing and Levy 1962) which is able to deal with complex solid-solutions 157 (Cannillo et al. 1983). Fully-ionized scattering factors of the appropriate chemical species were 158 used for non-tetrahedrally coordinated cation sites, whereas appropriate combinations of neutral vs. 159 ionized scattering-factors were used for the T sites and the O atoms (see Oberti et al. 1992). The 160 electron density in the A cavity was modelled using three different subsites, of which only A(m) and 161 A(2) are likely to be occupied by A-group cations. Site populations were derived and optimized by 162

combining refined site-scattering values with refined mean bond lengths, and using all the information derived from statistical treatment of the amphibole data base. More details on the procedure are given in the following sections. Refined atom coordinates and displacement parameters are given in Table 2; selected interatomic distances and angles are given in Table 3, and refined sitescattering values (Hawthorne et al. 1995) are listed in Table 4.

168

169 - Neutron diffraction

A crystal of K(1) (1.5 x 1.2 x 0.75 mm³), optically free of defects and twinning, was selected 170 under a transmitted-light polarizing microscope for the neutron-diffraction experiment. Neutron 171 172 Laue data were collected at room temperature on the Laue diffractometer KOALA on the OPAL reactor at the Australian Nuclear Science and Technology Organisation. KOALA, which is 173 174 essentially a clone of the Laue diffractometer VIVALDI at the Institut Laue-Langevin, Grenoble (McIntyre et al. 2006; Edwards 2011; Gatta et al. 2012), uses the Laue diffraction technique on an 175 176 unmonochromated thermal-neutron beam with a large solid-angle (8 sterad) cylindrical image-plate detector to increase the detected diffracted intensity by one to two orders of magnitude compared to 177 178 a conventional monochromatic experiment. The crystal was mounted with the high-symmetry axis well away from the single instrument-rotation axis, in order to avoid bias in the final refined 179 anisotropic-displacement parameters due to the blind region in reciprocal space around the rotation 180 181 axis. Twenty-nine Laue diffraction patterns, each accumulated over 60 min on average and with two different crystal orientations to ensure the best coverage of reciprocal space, were collected at 15° 182 intervals of rotation of the K(1) crystal. Two crystal orientations were used to ensure the best 183 coverage of reciprocal space and avoid bias in the final refined anisotropic-displacement parameters 184 due to the blind region in reciprocal space around the single rotation axis in one orientation. Data 185 extends to a minimum d spacing of ~ 0.5 Å (with $\lambda_{min} = 0.8$ Å and $\lambda_{max} = 5.2$ Å). 186

The Laue patterns were indexed on the basis of the unit-cell parameters previously measured by single-crystal X-ray diffraction (Table 1) and compatible with the reflection conditions of the space group C2/m, using the program LAUEGEN of the Daresbury Laboratory Laue Suite (Campbell 1995; Campbell et al. 1998), and the reflections were corrected for background and integrated using the program INTEGRATE+, which uses a two-dimensional version of the minimum $\sigma(I)/I$ algorithm (Wilkinson et al. 1988). Corrections for absorption were deemed unnecessary due to the small crystal volume and low fraction of highly absorbing elements. Reflections were normalised to the incident wavelength, using a curve derived by comparing equivalent reflections and multiple observations, and corrected for the different angles of incidence to the detector surface via the local program LAUE4 (Piltz 2012), and $|F_0| > 4\sigma(F_0)$. Diffraction peaks were indexed on the basis of the unit-cell parameters previously measured by single-crystal X-ray diffraction (Table 1) and compatible with the reflection conditions of the space group C2/m.

Anisotropic crystal-structure refinement based on the neutron intensity data was done in the 199 space group C2/m using the SHELX-97 software (Sheldrick 2008), starting from a structure model 200 without any H atom included. The neutron scattering lengths of Na, Ca, Fe, Ti, Al, Si, O and H 201 202 were taken from Sears (1986). Secondary isotropic extinction was corrected by the formalism of 203 Larson (1967), as implemented in the SHELXL-97 package. The structure refinement was done with the following strategy: the A(m) site was modelled with Na alone; the M(1), M(2), and M(3)204 sites were modelled as occupied by Fe and Ti (and their fractions were mutually refined) and M(4)205 by Ca and Na (fractions mutually refined); the T(1) and T(2) sites were modelled with Al and Si 206 207 (and their fractions mutually refined). Convergence was reached with one intense negative residual peak (at $x \sim 0.208$, $y \sim 0$, $z \sim 0.768$) in the final difference-Fourier map of the nuclear density. 208 209 Further cycles of refinement were done with H assigned to these coordinates, and its site occupancy was refined. Refinement of all variable parameters converged with all the principal mean-square 210 atomic displacement parameters positive, including those for the H site. The variance-covariance 211 212 matrix showed no significant correlation among the refined parameters. Site coordinates and displacement parameters are listed in Table 2; selected interatomic distances and angles are given in 213 Table 3, and refined site-scattering values and aggregate neutron-scattering lengths are listed in 214 Table 4. 215

216

217

- Microprobe analysis

Subsequent to the collection of the X-ray intensity data, the same fragment of the crystal K(1) was mounted in epoxy, polished, and analysed by electron- and ion-microprobe techniques (for H and Li; Li was found to be at the wt ppm level) following the procedures described by Oberti et al. (1992) and Ottolini et al. (1993, 2001). Chemical compositions (Table 5) are the mean of at least 10 determinations on the same crystal; because no significant compositional zoning was observed, they should be representative of the bulk composition of the sample. Formulae were calculated on the basis of 24(O,OH,F) = 24 *apfu*. The total number of cations was adjusted so as to obtain the best fit

between the refined group-site-scattering values and those calculated from EMP analysis. Given that 225 OH and F were measured by SIMS, this procedure implies a unique solution for the calculation of Fe³⁺ 226 within the uncertainty of the SIMS analysis. The initial cation assignment was done as discussed by 227 Hawthorne and Oberti (2007). This showed C-group cations (*i.e.*, Mg, Fe²⁺, Mn²⁺, Ni, Al, Fe³⁺, Cr³⁺, 228 Ti⁴⁺) in excess of 5.0 *apfu*, as is commonly the case in calcium amphiboles. What specific C-group 229 cations should be assigned to the B-group? Usually, one assigns Mn^{2+} and Fe^{2+} in that order, using the 230 site preference ${}^{M(4)}Mn^{2+} > {}^{M(4)}Fe^{2+}$ observed in monoclinic Fe-Mg-Mn amphiboles as the rationalization 231 for this procedure. In the present case, we know the total scattering from both the B-group $[\equiv M(4)$ site] 232 and the C-group $[\equiv M(1), M(2) \text{ and } M(3) \text{ sites}]$ cations, and can check the assignment on this basis. In 233 conclusion, all Mn^{2+} was assigned to M(4) together with 0.73 *apfu* of Fe²⁺. The resultant unit-formula 234 235 assignment is shown in Table 5.

236

237 Results: Site assignment of cations

238 - T-group cations

Initially, the Al contents at the T(1) and T(2) sites were allowed to vary during refinement of 239 the neutron data. However, the resultant site-populations were far too Al-rich for (1) the chemical 240 formula of the crystal; (2) the observed $\langle T(1)-O \rangle$ and $\langle T(2)-O \rangle$ distances, and (3) the grand $\langle T-O \rangle$ 241 distance. We suspected that too much scattering (all the cation sites) was being varied in the neutron 242 refinement, leading to poor scaling, and hence we refined the structure with the site populations of the 243 T(1) and (2) sites fixed at their values obtained from the X-ray structure refinement. Indeed, the < T(1)-244 O> and < T(2)-O> distances derived from the neutron refinement are almost identical to those obtained 245 from the X-ray refinement. According to the systematic work reported in Oberti et al. (2007), in 246 amphiboles the majority of ^TAl occurs at the T(1) site and only minor ^TAl occurs at the T(2) site. 247 Moreover, the $T^{(1)}Al$ content can be obtained from the regression equation of Oberti et al. (2007) $[T^{(1)}Al$ 248 = (< T(1)-O> - 1.6193) · 34.2199], which gives 1.84 *apfu*. Because the size of the T(2) tetrahedron 249 depends on many compositional factors (including the Fe content), the amount of $T^{(2)}Al$ cannot be 250 estimated quantitatively by a similar equation, and should be calculated by difference from the ^TAl 251 values obtained from EMP analysis: 1.97 - 1.84 = 0.13 apfu. Thus, T(1) = 2.16 Si + 1.84 Al and T(2) =252 3.87 Si + 0.13 Al apfu. 253

254

255 - C-group cations

The assignment of site populations to the M(1), M(2) and M(3) sites is the major goal of this 256 work. Inspection of Table 5 shows that there are five cations to be assigned to these sites: Mg^{2+} , 257 Fe^{2+} , Fe^{3+} , Al^{3+} and Ti^{4+} . We have assigned all Mn^{2+} and some Fe^{2+} to M(4), as indicated by the 258 chemical formula and the refined site-scattering values at M(4). With a combination of just two 259 different sets of diffraction data, we cannot uniquely determine site populations for five scattering 260 species over three sites (Hawthorne 1983), and hence we must reduce this number. In terms of 261 scattering, we may treat Fe^{2+} and Fe^{3+} as Fe, and assign ^CAl to M(2) in accord with the known site-262 preference for this species in amphiboles containing significant Fe (Hawthorne and Oberti 2007; 263 Oberti et al. 2007). This leaves three scattering species to be refined over three sites, M(1,2,3). 264

We may write the refined site-scattering values from the X-ray and neutron-scattering experiments as follows:

267 268

 $\begin{array}{ll}
269 & 12 \cdot {}^{Mg}a^{M(1)} + 26 \cdot {}^{Fe}a^{M(1)} + 22 \cdot {}^{Ti}a^{M(1)} = {}^{XRAY}S^{M(1)} \\
270 & 12 \cdot {}^{Mg}a^{M(2)} + 26 \cdot {}^{Fe}a^{M(2)} + 22 \cdot {}^{Ti}a^{M(2)} = {}^{XRAY}S^{M(2)} - 13 \cdot {}^{Al}a^{M(2)} \\
271 & 12 \cdot {}^{Mg}a^{M(3)} + 26 \cdot {}^{Fe}a^{M(3)} + 22 \cdot {}^{Ti}a^{M(3)} = {}^{XRAY}S^{M(3)} \\
272 \\
273 & {}^{Mg}B \cdot {}^{Mg}a^{M(1)} + {}^{Fe}B \cdot {}^{Fe}a^{M(1)} + {}^{Ti}B \cdot {}^{Ti}a^{M(1)} = {}^{NEUTRON}S^{M(1)} \\
274 & {}^{Mg}B \cdot {}^{Mg}a^{M(2)} + {}^{Fe}B \cdot {}^{Fe}a^{M(2)} + {}^{Ti}B \cdot {}^{Ti}a^{M(2)} = {}^{NEUTRON}S^{M(2)} - {}^{Al}B \cdot {}^{Al}a^{M(2)} \\
275 & {}^{Mg}B \cdot {}^{Mg}a^{M(3)} + {}^{Fe}B \cdot {}^{Fe}a^{M(3)} + {}^{Ti}B \cdot {}^{Ti}a^{M(3)} = {}^{NEUTRON}S^{M(3)} \\
276 \\
276 & {}^{MUUUUUUUU}
\end{array}$

where ^{XRAY}S^{M(1)} is the refined site-scattering value for the M(1) site from the refinement of the Xray data, ^{NEUTRON}S^{M(1)} is the refined site-scattering value for the M(1) site from the refinement of the neutron data, ^{Mg}B is the neutron scattering length for Mg, ^{Mg}a^{M(1)} is the amount (site population) of Mg assigned to the M(1) site (etc. for the other sites and the other scattering species). Stoichiometry requires the following equations (*i.e.*, there are no vacancies at the M(1,2,3) sites):

282

283 $M_{g}a^{M(1)} + F_{e}a^{M(1)} + T_{i}a^{M(1)} = 2$ 284 $M_{g}a^{M(2)} + F_{e}a^{M(2)} + T_{i}a^{M(2)} = 2 - A_{a}a^{M(2)}$

- 285 $Mg_a M^{(3)} + Fe_a M^{(3)} + Ti_a M^{(3)} = 1$
- 286

There are nine site-population variables $Mga^{M(1)}$ etc. and nine equations, and we may solve for the 287 site populations ${}^{Xa^{M(n)}}$ where X = Mg, Fe and Ti, and n = 1-3. The resultant values are given in 288 Table 6. They confirm that Ti⁴⁺ occurs in all the three octahedrally coordinated sites in this sample 289 (as suggested by Tiepolo et al. 1999). The unit-cell contents derived from the chemical analysis 290 (Table 5) provide a check on the accuracy of these site populations by comparing the sums of the 291 different scattering species assigned from the diffraction experiments with the same scattering 292 species determined by electron-microprobe analysis, as is done in the two columns to the right in 293 Table 6. The closeness of the two sets of compositions strongly suggests that the derived site 294 295 populations are correct.

The values given in Table 6 are derived solely from the site-scattering values and do not take into account the presence of both Fe^{2+} and Fe^{3+} . We may examine this issue in terms of the geometrical parameters (mean bond-lengths and distortion) at the M(1), M(2) and M(3) sites, in two ways as we now describe:

1) Use optimised ideal mean bond-lengths for complete occupancy of the M(1,2,3) sites by Mg 300 (2.078 Å), Fe²⁺ (2.125 Å), Fe³⁺ (2.025 Å), Ti⁴⁺ (1.960 Å) and Al (1.929 Å) for (OH, F, Cl) 301 amphiboles based on the CNR-IGG database in Pavia (Oberti et al. 2007), adjusted for the effect of 302 F at O(3) on the $\langle M(1) - O \rangle$ and $\langle M(3) - O \rangle$ distances (-0.13 F *apfu*), and taking into account the 303 observation that significant Fe^{3+} content at the M(1) site causes a considerable increase in the M(1)304 octahedron distortion parameter Δ (Brown and Shannon 1973) from the common value of 2-3. The 305 value of Δ for the M(1) octahedron in crystal K(1) is 2.49, thus excluding the presence of Fe³⁺ at 306 M(1). With no Fe³⁺ at M(1), the calculated $\langle M(1) - O \rangle$ distance is 2.064 Å, already shorter than the 307 refined value of 2.078 Å. Thus Fe^{3+} is either ordered at the M(2) site, which would lead to a 308 calculated $\langle M(2)-O \rangle$ distance of 2.049 Å and a calculated $\langle M(3)-O \rangle$ distance of 2.069 Å (to be 309 compared to the refined values of 2.074 Å and 2.070 Å, respectively), or is distributed between the 310 M(2) and the M(3) sites. This latter hypothesis would make the overall shrinking of the strip of 311 312 octahedra more homogeneous in crystal K(1), which is consistent with the presence of a significant oxo component. This issue will be explored further in the last section of this paper. 313

2) The regression equations of Hawthorne and Oberti (2007, Table 7) may be used to calculate the Fe³⁺ content of each M(1,2,3) site using the occupancies given in Table 6. The equation for the M(2)octahedron involves just the aggregate radius of the constituent M(2) cations, whereas the equations for the M(1) and M(3) octahedra take into account not only the aggregate radius of the constituent 318 M(1) and M(3) cations, but also the aggregate radius of the anion(s) at O(3) and the cations at M(4), 319 the amount of ^TAl and the aggregate net charge at the *A* site. The results are shown below:

320

321
$$M(1)$$
 Fe³⁺ = 0.00 apfu, $< M(1)$ -O>_{obs} = 2.078 and $< M(1)$ -O>_{calc} = 2.077 Å;

322
$$M(2)$$
 Fe³⁺ = 0.00 apfu, $\langle M(2) - O \rangle_{obs} = 2.073$ and $\langle M(2) - O \rangle_{calc} = 2.073$ Å;

323
$$M(3)$$
 Fe³⁺ = 0.10 apfu, $\langle M(3) - O \rangle_{obs} = 2.070$ and $\langle M(3) - O \rangle_{calc} = 2.070$ Å.

324

This gives a significantly lower Fe³⁺ content than that obtained by formula recalculation with the constrain of H and F measured by SIMS (*i.e.*, 0.33 *apfu*). The predicted $\langle M$ -O> values have standard errors of estimate of 0.0027, 0.0053 and 0.0058 Å for M((1), M(2) and M(3), respectively,which at their limits would predict a further Fe³⁺ content of 0.023 Fe³⁺. Anyway, these results indicate that Fe³⁺ is strongly ordered at the M(3) site and is in reasonable accord with electroneutrality using the neutron-derived H content of 1.02(2) *apfu*.

The presence of a significant amount of Ti^{4+} at the M(1) site is usually detected by a much 331 higher value of the anisotropic displacement parameters (*adp*) refined at this site than at the M(2)332 and M(3) sites, indicating a shift of the Ti⁴⁺ ions toward the edge of the octahedron connecting the 333 two O(3) sites (Tiepolo et al. 1999). In some cases, two split positions, M(1) and M(1'), can even be 334 refined (e.g., Hawthorne et al. 2000). In this work, a higher value of the adp at M(1) and M(3) was 335 obtained from the X-ray data but not from the neutron data. The presence of Ti^{4+} at M(1), M(2) and 336 M(3), as well as the presence of significant F, may be the reason for the unusual distortion pattern of 337 the ribbon of octahedra observed in crystal K(1), which is generally more shrunk than expected 338 based on the ideal <cat-O> distances used in the crystal-chemical analysis of amphiboles (cf. Oberti 339 et al. 2007 and the discussion above). 340

341

342

B-group cations

Although two different subsites (*i.e.*, M(4) and M(4'); Table 2) can be refined from X-ray data due to the oblong shape of the electron density, their distance is close to the resolution of the data. Hence, the total scattering obtained is more accurate than that based on the refinement of a unique position, but its partitioning cannot be considered accurate. However, inspection of the cat-O distances and comparison with the shape of the electron density found in many other amphibole compositions show that smaller cations (Mg, Fe²⁺, Mn²⁺) order at the M(4') position (Oberti et al. 2007). Cations were assigned to the B-group in the usual fashion (*e.g.*, Hawthorne et al. 2012). Those C-group cations in excess of 5 *apfu* were assigned as B-group cations in the following sequence: $Mn^{2+}>Fe^{2+}>Mg$, and all Ca was assigned to this site until the site is completely occupied (Hawthorne et al. 2012). The resultant calculated X-ray/neutron site-scattering values (Table 4) are in close agreement with the analogous refined values for both the X-ray and neutron refinements.

354

355 - A-group cations

The site populations at the A(m), A(2) and A sites were refined unconstrained in both X-ray and neutron refinements. Although the model used for neutron data includes only the A(m) subsite, as the A and A(2) site populations were found not to be significant, the aggregate refined scattering/neutron-scattering value in the A cavity is close to that calculated from the A-group cations based on electron- and ion-microprobe analysis (Table 4). The higher value obtained by Xrays structure refinement may be due to the relatively high values of the atomic displacement parameters (Table 2).

363

364 - W-group anions

The site population of the H site was refined directly in the neutron structure refinement, with convergence to the following value: 1.02(2) H + $0.98 \square apfu$. This may be compared to the value derived by SIMS of 0.87 H *apfu*. Ottolini and Hawthorne (2001) give an uncertainty of 5-10% for the determination of H by SIMS for the analytical method used here, *i.e.*, 0.04-0.09 *apfu* for crystal K(1). The upper limit of the uncertainty plus the standard deviation of the H value derived by neutron scattering suggest that there is little significant difference between these two values.

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Discussion and conclusions

374 - Hydrogen bonding

The neutron structure refinement provides a clear picture of the H site location, anisotropic displacement regime and potential H-bonding configuration. The coordinates of the H site obtained in this study are consistent with previous experimental findings based on single-crystal neutron structure refinement reported by Pechar et al. (1989) for a kaersutite from Bohemia. The refined O-H bond distance of this study is 0.951(2) Å (Table 3), whereas that reported by Pechar et al. (1989)

is 0.93(1) Å. A correction for the "riding motion effect" (following the protocol of Busing and Levy 380 1964) gives 0.982 Å (Table 3), showing that the O-H distance is significantly affected by the 381 motion of the proton. The magnitude and orientation of the displacement ellipsoid of the H site of 382 this study is shown in Fig. 2, along with the potential H-bonding network. The shape of the 383 ellipsoid (root-mean-square components: R1:R2:R3 = 1.89:1.84:1, with R1 and R2 dispersed on 384 (100), Fig. 2) is strongly influenced by the geometry of the weak interactions with the oxygens of 385 the two independent tetrahedra. Inspection of the difference-Fourier map of the nuclear density does 386 not show any evidence of splitting of the H site. Pechar et al. (1989) did not report any indication of 387 possible hydrogen bonding. However, our data lead us to consider that at least three potential weak 388 H-bonds may occur: with $O(6) \ge 2$ and O(7) as acceptors (i.e., H···O(6)= 2.763(2) Å, O(3)···O(6) 389 =3.316(1) Å and O(3)-H···O(6) $=117.86(1)^{\circ}$; H···O(7)= 2.694(3) Å, O(3)···O(7) = 3.270(2) Å and 390 O(3)-H···O(7) = 119.6(2)°, Table 3). A similar trifurcated configuration is observed in trioctahedral 391 phyllosilicates, in which the *donor* is an oxygen site shared by three adjacent octahedra (belonging 392 to the octahedral sheet) and the acceptors are oxygens of the superimposed 6-membered ring of 393 tetrahedra (belonging to the tetrahedral sheet) (e.g., Gatta et al. 2011, 2013, 2014). In all these 394 cases, the Odonor-H···Oacceptor angles range between 115-140°. 395

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- 397

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The relation between ^[6]Ti⁴⁺ and ^{O(3)}O²⁻

The Mg-Fe-Ti site populations of the C cations calculated from the refined X-ray and 398 neutron site-scattering values are in excellent agreement with the analysed chemical compositions 399 (Table 6). Ti⁴⁺ is preferentially ordered at the M(1) site, in accord with the dominance of 400 substitution (2) (see above). The refined Ti⁴⁺ content of M(1) is 0.38 apfu (Table 6); using the 401 neutron site population for hydrogen at the H site, the site population of O(3) is 1.02 (OH)⁻ + 0.20 F⁻ 402 + 0.78 O²⁻. These values, $M^{(1)}Ti^{4+} = 0.38$ apfu and $O^{(3)}O^{2-} = 0.78$ apfu, are in almost exact accord 403 with substitution (2), suggesting that the incorporation of O^{2-} at O(3) is controlled by the occurrence 404 of Ti^{4+} at M(1), and the bond-valence requirements of Ti^{4+} at M(1) are met primarily by very short 405 bonds to O^{2-} at the O(3) site. 406

However, there is significant Ti^{4+} at the M(2) and M(3) sites (Table 6). Incorporation of Ti^{4+} at the M(2) site is a result of substitution (3) and is unrelated to dehydrogenation in amphiboles. The incorporation of Ti^{4+} at M(3) is more problematic. Incorporation of Ti^{4+} at M(3) could in principle be related to dehydrogenation as the M(3) site is also coordinated by two O(3) sites. However, in

amphibole K(1), all O²⁻ at O(3) is associated with Ti⁴⁺ at M(1), and Ti⁴⁺ at M(3) cannot be 411 associated independently with O^{2-} at O(3) as all O^{2-} at O(3) is already associated with Ti^{4+} at M(1). 412 This being the case, it is not clear how the bond-valence requirements of Ti^{4+} at M(3) can be 413 satisfied. Possibly, M(3)Ti⁴⁺ could be incorporated into the amphibole structure by, for example, a 414 substitution such as ${}^{M(3)}\text{Ti}^{4+} + 2 {}^{M(2)}\text{Mg}^{2+} \rightarrow {}^{M(3)}\text{Mg}^{2+} + 2\text{Al}^{3+}$, that is locally spatially associated with 415 substitution (2), such that ${}^{M(3)}$ Ti⁴⁺ bonds to ${}^{O(3)}O^{2-}$. An analogous mechanism could also account for the 416 occurrence of Fe^{3+} at M(3). 417

The Ti⁴⁺ content at the M(1) site compensates for 0.76 apfu O²⁻ at O(3). This leaves a 418 difference of 0.02 (0.78 - 0.76 from the neutron H-site occupancy) to 0.17 apfu (0.93 - 0.76 from the neutron H-site occupancy)419 the SIMS H content) to be compensated by additional substitutions. Given the chemical formula of 420 the sample, two possibilities remain: (1) the additional positive charge is provided by half the Ti⁴⁺ 421 occurring at M(3), which implies substitution (6) below: 422

423
$$M^{(3)}\text{Ti}^{4+} + 2^{O(3)}\text{O}^{2-} \rightarrow M^{(3)}\text{Mg}^{2+} + 2^{O(3)}(\text{OH})^{-}$$
 (6),

or (2) the additional positive charge is provided by up to 0.17 Fe^{3+} apfu occurring at M(3) according 424 to substitution (7): 425

$${}^{M(3)}\mathrm{Fe}^{3+} + {}^{\mathrm{O}(3)}\mathrm{O}^{2-} \rightarrow {}^{M(3)}\mathrm{Fe}^{2+} + {}^{\mathrm{O}(3)}(\mathrm{OH})^{-}$$
 (7).

Note that substitution (7) is common in oxo-amphiboles where the Ti content is low or 427 dehydrogenation has occurred after crystallization (e.g., Oberti et al. 2016). 428

Above, we considered two approaches concerning the amount and distribution of Fe^{3+} at the 429 M(1), M(2) and M(3) sites. Both exclude the presence of Fe³⁺ at the M(1) site and are compatible 430 with the occurrence of Fe^{3+} at the M(3) site. In this regard, crystal K(1) is unusual as 431 dehydrogenation of Fe-rich amphiboles generally proceeds according to substitution (8) below: 432

433

$${}^{M(1)}\mathrm{Fe}^{3+} + {}^{\mathrm{O}(3)}\mathrm{O}^{2-} \rightarrow {}^{M(1)}\mathrm{Fe}^{2+} + {}^{\mathrm{O}(3)}(\mathrm{OH})^{-}$$
(8)

The second approach predicts 0.13 apfu Fe³⁺ at M(3), within the range 0.02-0.17 apfu required by 434 electroneutrality, highlighting the unusual nature of this amphibole. 435

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444	References
445	Brown ID (1981) The bond-valence method: an empirical approach to chemical structure
446	and bonding. In M. O'Keeffe and A. Navrotsky, Eds., Structure and Bonding in Crystals, Vol. 2,
447	1–30. Academic Press, New York.
448	Brown ID (2002) The chemical bond in inorganic chemistry. The Bond Valence Model.
449	Oxford University Press, U.K.
450	Brown ID, Shannon RD (1973) Empirical bond strength-bond lengths curves for oxides.
451	Acta Crystallogr A29:266-282.
452	Busing WR, Martin KO, Levy HA (1962) ORFLS. Report ORNL-Tm-305 Oak Ridge Natl.
453	Lab. Oak Ridge, Tn.
454	Busing WR, Levy HA (1964) The effect of thermal motion on the estimation of bond lengths
455	from diffraction measurements. Acta Crystallogr 17:142-146.
456	Campbell JW (1995) LAUEGEN, an X-windows-based program for the processing of Laue
457	diffraction data. J Appl Crystallogr 28:228-236.
458	Campbell JW, Hao Q, Harding MM, Nguti ND, Wilkinson C (1998) LAUEGEN version 6.0
459	and INTLDM. J Appl Crystallogr 31:496-502.
460	Cannillo E, Germani G, Mazzi F (1983) New crystallographic software for Philips
461	PW11000 single crystal diffractometer. CNR Centro di Studio per la Cristallografia, Internal Report
462	2.
463	Della Ventura G, Robert JL, Bény JM (1991) Tetrahedrally coordinated Ti ⁴⁺ in synthetic Ti-
464	rich potassic richterites: evidence from XRD, FTIR and Raman study. Am Mineral 76:1134-1140.
465	Della Ventura G, Robert JL, Bény JM, Raudsepp M, Hawthorne FC (1993) The OH-F
466	substitution in Ti-rich potassium-richterites: Rietveld structure refinement and FTIR and micro-Raman
467	spectroscopic studies of synthetic amphiboles in the system K2O-Na2O-CaO-MgO-SiO2-TiO2-H2O-
468	HF. Am Mineral 78:980-987.
469	Della Ventura G, Robert JL, Hawthorne FC, Prost R (1996) Short-range disorder of Si and Ti
470	in the tetrahedral double-chain unit of synthetic Ti-bearing potassium-richterite. Am Mineral 81: 56-
471	60.

472	Della Ventura G, Oberti R., Hawthorne FC, Bellatreccia F (2007) FTIR spectroscopy of Ti-rich
473	pargasites from Lherz and the detection of O ^{2B} at the anionic O3 site in amphiboles. Am Mineral 92,
474	1645-1651.
475	Edwards AJ (2011) Neutron Diffraction: Recent Applications to Chemical Structure
476	Determination. Austral J Chem 64:869–872.
477	Gatta GD, McIntyre GJ, Sassi R, Rotiroti N, Pavese A (2011) Hydrogen-bond and cation
478	partitioning in 2M1-muscovite: A single-crystal neutron-diffraction study at 295 and 20 K. Am Mineral
479	96:34-41.
480	Gatta GD, McIntyre GJ, Bromiley G, Guastoni A, Nestola F (2012) A single-crystal neutron
481	diffraction study of hambergite, Be2BO3(OH,F). Am Mineral 97:1891-1897.
482	Gatta GD, Merlini M, Valdrè G, Liermann H-P, Nénert G, Rothkirch A, Kahlenberg V, Pavese
483	A (2013) On the crystal structure and compressional behaviour of talc: a mineral of interest in
484	petrology and material science. Phys Chem Minerals 40:145-156.
485	Gatta GD, Nénert G, Guastella G, Lotti P, Guastoni A, Rizzato S (2014) A single-crystal
486	neutron and X-ray diffraction study of a Li,Be-bearing brittle mica. Min Mag 78:55–72.
487	Hawthorne FC (1983) Quantitative characterization of site occupancies in minerals. Am
488	Mineral 68:287-306.
489	Hawthorne FC (1997) Short-range order in amphiboles: A bond-valence approach. Can
490	Mineral 35:201–216.
491	Hawthorne FC, Ungaretti L, Oberti R (1995) Site populations in minerals: terminology and
492	presentation of results of crystal-structure refinement. Can Mineral 33:907-911.
493	Hawthorne FC, Oberti R, Zanetti A, Czamanske GK (1998) The role of Ti in hydrogen-
494	deficient amphiboles: Sodic-calcic and sodic amphiboles from Coyote Peak, California. Can
495	Mineral 36:1253-1265.
496	Hawthorne FC, Cooper MA, Grice JD, Ottolini L (2000) A new anhydrous amphibole from
497	the Eifel region, Germany: Description and crystal structure of obertiite,
498	NaNa2(Mg3Fe ³⁺ Ti ⁴⁺)Si8O22O2. Am Mineral 85: 236–241.
499	Hawthorne FC, Oberti R (2007) Amphiboles: Crystal chemistry. Rev Mineral Geochem 67:1-
500	54.
501	Hawthorne FC, Oberti R, Harlow GE, Maresch W, Martin RF, Schumacher JC, Welch MD
502	(2012) Nomenclature of the amphibole super-group. Am Mineral 97:2031-2048.

503 Kitamura M, Tokonami M, Morimoto N (1975) Distribution of Ti in oxy-kaersutite. Contr 504 Mineral Petrol 51:167-172.

505 Konzett J (1997) Phase relations and chemistry of Ti-rich K-richterite-bearing mantle 506 assemblages: an experimental study to 8 GPa in a Ti-KNCMASH system. Contrib Mineral Petrol 507 128:385-404.

508 Konzett J, Sweeney RJ, Thompson AB, Ulmer P (1997) Potassium amphibole stability in 509 the upper mantle: an experimental study in a peralkaline KNCMASH system to 8.5 GPa. J Petrol 510 38: 537-568.

511 Larson AC (1967) Inclusion of secondary extinction in least-squares calculations. Acta
512 Crystallogr 23:664 – 665.

513 Leake BE (1968) A catalog of analysed calciferous and subcalciferous amphiboles together 514 with their nomenclature and associated minerals. Geol Soc Am Spec Paper 98, 210p.

- McIntyre GJ, Lemée-Cailleau MH, Wilkinson C (2006) High-Speed Neutron Laue Diffraction
 Comes of Age. Physica B 385-386:1055-1058.
- 517 Oberti R, Ungaretti L, Cannillo E, Hawthorne FC (1992) The behaviour of Ti in amphiboles: I. 518 Four- and six-coordinated Ti in richterite. Eur J Mineral 4:425-439.
- 519 Oberti R, Hawthorne FC, Cannillo E, Cámara F (2007) Long-range order in amphiboles.
 520 Rev Mineral Geochem 67:125-171.
- 521 Oberti R, Boiocchi M, Zema M, Della Ventura G (2016) Synthetic potassic-ferro-richterite:

522 1. Composition, crystal structure refinement and HT behavior by *in operando* single-crystal X-ray

- 523 diffraction. Can Mineral (*in press*).
- 524 Ottolini L, Hawthorne FC (2001) SIMS ionization of hydrogen in silicates: A case study of 525 kornerupine. J Anal At Spectrom 16:1266–1270.
- 526 Ottolini L, Bottazzi P, Vannucci R (1993) Quantification of Li, Be and B in silicates by 527 secondary ion mass spectrometry using conventional energy filtering. Anal Chem 65:1960–1968.
- Pechar F, Fuess H, Joswig W (1989) Refinement of the crystal structure of kaersutite
 (Vlcíhora, Bohemia) from neutron diffraction. Neues Jahrb Mineral Monat 89:137-143.
- Piltz RO (2011) Accurate data analysis for the Koala and VIVALDI neutron Laue
 diffractometers. Abstracts of the XXII IUCr Congress, Madrid (Spain) 22-30 August 2011. Acta
 Crystallogr A67:C155.

Saxena SK, Ekström TK (1970) Statistical chemistry of calcic amphibole. Contrib Mineral
Petrol 26:276-284.

Sears, V.F. (1986) Neutron Scattering Lengths and Cross-Sections. In K. Sköld and D.L.
Price, Eds., Neutron Scattering, Methods of Experimental Physics, Vol. 23A, 521-550. Academic
Press, New York.

538 Sheldrick, G.M. (2008) A short history of SHELX. Acta Crystallogr A64:112-122.

Tiepolo M, Zanetti A, Oberti R (1999) Detection, crystal-chemical mechanisms and petrological implications of ^[6]Ti⁴⁺ partitioning in pargasite and kaersutite. Eur J Mineral 11:345-354.

- 542 Wilkinson C, Khamis HW, Stansfield RFD, McIntyre GJ (1988) Integration of single-crystal
 543 reflections using area multidetectors. J Appl Crystallogr 21:471-478.
- Wilshire HG, Meyer CE, Nakata JK, Calk LC, Shervais JW, Nielson JE, Schwarzman EC
 (1988) Mafic and ultramafic xenoliths from volcanic rocks of the Western United States. U.S.
 Geological Survey Professional Paper, 1443, 179 pp.

Table 1. Summary of crystallographic, data collection and structure
refinement information.

	K(1) X-ray	K(1) neutron [*]
a (Å)	9.882(4)	9.882(4)
<i>b</i> (Å)	18.084(7)	18.084(7)
<i>c</i> (Å)	5.312(1)	5.312(1)
β (°)	105.26(6)	105.26(6)
$V(Å^3)$	915.84	915.84
Space group	C2/m	<i>C</i> 2/ <i>m</i>
Ζ	2	2
Crystal size (mm)	0.75 x 0.43 x 0.22	2.0 x 1.5 x 0.9
Rad./mono.	Μο <i>Κ</i> α	Thermal neutrons / Laue, $(\lambda_{min}=0.8, \lambda_{max}=5.2 \text{ Å})$
No. of reflections	16801	31909

1254

0.7

3.70

129

1.78

2.28

3099

2261

0.5

6.13

116

5.25

7.45

+2.1/-2.8 fm/Å3

Unique reflections

No. of refined parameters

Maximum residual peaks

 d_{\min} (Å)

 $R_{\rm sym}\%$

 $R_{\rm all}$ %

** $R_1(F)$ %

Unique observed reflections 1210

*Cell dimensions taken from X-ray data; ** $R_1(F)$ % calculated on the basis of $F_o > 3\sigma(F_o)$ for the X-ray data and $F_o > 4\sigma(F_o)$ for the neutron data; ***close to the M(1) site, can be attributed to the ordering of Ti⁴⁺.

***0.47 e-/Å3

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			1	1		1				
Site	x	У	z	U^{11*}	U^{22*}	U^{33*}	U^{12*}	U^{13*}	U^{23*}	$U^{ m eq^*}$
K(1) –	X-ray									
<i>T</i> (1)	0.28234(5)	0.08503(3)	0.30227(10)	0.0058(2)	0.0066(2)	0.0066(2)	-0.0009(2)	0.0007(2)	-0.0003(2)	0.0065(2)
<i>T</i> (2)	0.29067(5)	0.17235(3)	0.81009(9)	0.0056(2)	0.0075(2)	0.0060(2)	-0.0009(2)	0.0011(2)	0.0001(2)	0.0064(2)
<i>M</i> (1)	0	0.08646(4)	1/2	0.0098(3)	0.0176(3)	0.0081(3)	0	0.0040(2)	0	0.0115(2)
<i>M</i> (2)	0	0.17730(4)	0	0.0065(4)	0.0073(3)	0.0067(4)	0	0.0018(2)	0	0.0068(2)
<i>M</i> (3)	0	0	0	0.0102(5)	0.0073(5)	0.0086(5)	0	-0.0009(3)	0	0.0093(3)
<i>M</i> (4)	0	0.27936(3)	1/2	0.0115(3)	0.0078(3)	0.0113(3)	0	0.0068(2)	0	0.0095(2)
<i>M</i> (4')	0	0.2621(3)	1/2	0	0	0	0	0	0	0.0171(10)
Α	0	1/2	0	0.076(7)	0.021(4)	0.076(7)	0	0.074(7)	0	0.048(4)
A(m)	0.0537(7)	1/2	0.1067(12)	0.053(3)	0.025(2)	0.047(3)	0	0.036(3)	0	0.037(2)
A(2)	0	0.4737(5)	0	0.051(5)	0.037(4)	0.079(7)	0	0.055(5)	0	0.049(4)
O(1)	0.10776(13)	0.08645(7)	0.2191(3)	0.0078(6)	0.0128(7)	0.0093(6)	-0.0012(4)	0.0017(5)	-0.0001(5)	0.0101(4)
O(2)	0.11914(13)	0.17219(7)	0.7283(3)	0.0073(6)	0.0114(5)	0.0098(6)	-0.0004(4)	0.0014(4)	0.0009(5)	0.0097(3)
O(3)	0.1082(2)	0	0.7150(4)	0.0092(8)	0.0156(8)	0.0136(9)	0	0.0021(7)	0	0.0130(5)
O(4)	0.3656(2)	0.24973(8)	0.7864(3)	0.0140(6)	0.0109(7)	0.0120(6)	-0.0027(5)	0.0037(5)	0.0008(5)	0.0123(4)
O(5)	0.34973(14)	0.13938(8)	0.1091(3)	0.0094(6)	0.0167(7)	0.0130(6)	-0.0005(5)	0.0012(5)	0.0051(5)	0.0134(4)
O(6)	0.34609(14)	0.11660(8)	0.6080(3)	0.0094(6)	0.0167(7)	0.0150(6)	0.0009(5)	0.0027(5)	-0.0051(5)	0.0138(4)
O(7)	0.3428(2)	0	0.2782(4)	0.0119(9)	0.0157(10)	0.0206(11)	0	0.0015(8)	0	0.0166(6)
$\overline{K(1)}$ -l	Veutron									
<i>T</i> (1)	0.28216(6)	0.08500(3)	0.30226(12)	0.0052(2)	0.0060(2)	0.0060(2)	-0.0006(1)	0.0006(2)	-0.0006(1)	0.0059(2)
<i>T</i> (2)	0.29056(5)	0.17238(3)	0.81021(11)	0.0059(2)	0.0065(2)	0.0060(2)	-0.0009(1)	0.0012(2)	0.0003(1)	0.0062(1)
<i>M</i> (1)	0	0.09012(3)	1/2	0.0074(3)	0.0023(2)	0.0059(3)	0	0.0016(2)	0	0.0052(2)
<i>M</i> (2)	0	0.17697(3)	0	0.0071(3)	0.0069(2)	0.0068(3)	0	0.0018(2)	0	0.0069(2)

Table 2. Atom coordinates and displacement parameters $(Å^2)$ for the amphibole of this work.

<i>M</i> (3)	0	0	0	0.0070(3)	0.0062(3)	0.0056(4)	0	0.0016(3)	0	0.0063(2)
<i>M</i> (4)	0	0.27828(4)	1/2	0.0119(3)	0.0145(3)	0.0112(3)	0	0.0065(3)	0	0.0119(2)
A(m)	0.0369(7)	1/2	0.0762(13)	0.089(7)	0.118(5)	0.093(6)	0	0.077(5)	0	0.090(3)
O(1)	0.10791(4)	0.08643(2)	0.21876(9)	0.0074(2)	0.0119(2)	0.0095(2)	-0.0012(1)	0.0012(1)	-0.0001(1)	0.0098(1)
O(2)	0.11925(4)	0.17211(2)	0.72815(8)	0.0066(2)	0.0104(1)	0.0095(2)	-0.0002(1)	0.0008(1)	0.0012(1)	0.0091(1)
O(3)	0.10828(6)	0	0.71491(13)	0.0095(2)	0.0142(2)	0.0148(3)	0	0.0013(2)	0	0.0132(1)
O(4)	0.36545(5)	0.24984(2)	0.78717(9)	0.0136(2)	0.0093(2)	0.0121(2)	-0.0029(1)	0.0035(2)	0.0001(1)	0.0116(1)
O(5)	0.34970(4)	0.13944(2)	0.10881(9)	0.0099(2)	0.0157(2)	0.0118(2)	-0.0009(1)	0.0010(2)	0.0050(1)	0.0128(1)
O(6)	0.34592(4)	0.11669(2)	0.60790(9)	0.0106(2)	0.0151(2)	0.0140(2)	0.0005(1)	0.0027(2)	-0.0057(1)	0.0134(1)
O(7)	0.34272(7)	0	0.27894(14)	0.0126(3)	0.0139(2)	0.0210(3)	0	0.0020(2)	0	0.0163(1)
Н	0.2080(3)	0	0.7672(7)	0.013(1)	0.044(2)	0.045(2)	0	0.004(1)	0	0.034(1)

* for K(1) - X-ray, the U^{ij} values are calculated from the refined β^{ij} values. The atom- displacement parameter takes the form: exp $[-2\pi^2 (h^2 a^{*2} U^{11} + k^2 b^{*2} U^{22} + l^2 c^{*2} U^{33} + 2 hka^* b^* U^{12} + 2 hla^* c^* U^{13} + 2 klb^* c^* U^{23})].$

	X-ray	Neutron		X-ray	Neutron
T(1) - O(1)	1.665(2)	1.661(1)	T(2) - O(2)	1.635(2)	1.633(1)
<i>T</i> (1)–O(5)	1.680(2)	1.683(1)	T(2)-O(4)	1.604(2)	1.604(1)
<i>T</i> (1)–O(6)	1.680(2)	1.680(1)	T(2)-O(5)	1.653(2)	1.651(1)
<i>T</i> (1)–O(7)	1.667(1)	1.666(1)	T(2)-O(6)	1.667(2)	1.667(1)
< <i>T</i> (1)–O>	1.673	1.673	< <i>T</i> (2)–O>	1.640	1.639
M(1)–O(1) x2	2.049(1)	2.052(1)	<i>M</i> (2)–O(1) x2	2.130(2)	2.125(1)
<i>M</i> (1)–O(2) x2	2.123(2)	2.075(1)	<i>M</i> (2)–O(2) x2	2.091(1)	2.092(1)
<i>M</i> (1)–O(3)	2.062(1)	2.112(1)	<i>M</i> (2)–O(4) x2	2.000(2)	2.002(1)
< <i>M</i> (1)–O>	2.078	2.080	< <i>M</i> (2)–O>	2.073	2.073
<i>M</i> (3)–O(1) x4	2.069(1)	2.068(1)	A-O(5) x4	3.058(2)	
M(3) - O(3) x2	2.071(2)	2.072(1)	A-O(6) x4	3.071(2)	
< <i>M</i> (3)–O>	2.070	2.070	A-O(7) x2	2.410(2)	
			<a-o></a-o>	2.846	
A(m) - O(5) x2	3.026(2)	3.027(4)			
A(m) - O(6) x2	2.633(2)	2.754(5)	M(4)–O(2) x2	2.421(2)	2.407(1)
A(m)–O(7)	2.485(2)	2.435(7)	M(4)–O(4) x2	2.328(2)	2.328(1)
A(m)–O(7)	2.516(2)	2.481(6)	<i>M</i> (4)–O(5) x2	2.656(2)	2.667(1)
< <i>A</i> (<i>m</i>)–O>	2.720	2.746	M(4)–O(6) x2	2.577(2)	2.591(1)
			< <i>M</i> (4)–O>	2.495	2.523
A(2) - O(5) x2	2.679(2)				
A(2)–O(6) x2	2.766(2)		<i>M</i> (4')–O(2) x2	2.180(1)	
A(2)–O(7) x2	2.456(2)		<i>M</i> (4')–O(4) x2	2.278(2)	
<a(2)–o></a(2)–o>	2.634	_	<i>M</i> (4')–O(5) x2	2.840(2)	
			<i>M</i> (4')–O(6) x2	2.811(2)	
			< <i>M</i> (4')–O>	2.527	_
O(3)–H		0.951(2)	O(3)–H…O(6) x2		117.86(1)
O(3)–H*		0.982	H…O(7)		2.694(3)
H…O(6)		2.763(2)	O(3)····O(7)		3.270(2)
O(3)····O(6)		3.316(1)	O(3)–H O(7)		119.6(2)

Table 3. Selected interatomic distances (Å) and angles (°) for the amphibole of this work.

	X-ray	(epfu)	Neutron (barns pfu	
Site	Refined	EMPA	Refined	
<i>M</i> (1)	34.08		9.144	
M(2)	30.83		10.487	
<i>M</i> (3)	17.62		5.120	
Sum C	82.53	82.62	24.75	
<i>M</i> (4)	36.45			
<i>M</i> (4')	4.14			
Sum B	40.59	40.52	9.53	
A(m)	6.14			
A(2)	4.26			
A	2.55			
Sum A	12.96	12.64	3.69	

Table 4.	Site-scatter	ing values	refined for	or the am	ohibole
of this w	ork.	-		_	

Table 5. Chemical composition (wt.%) and unit formula (*apfu*) derived from EMP and SIMS analysis (H, F) of crystal K(1). *bdl* = below detection limit.

SiO ₂	39.99(47)
TiO ₂	6.07(11)
Al ₂ O ₃	12.02(24)
Cr ₂ O ₃	0.02(2)
*Fe ₂ O ₃	2.88
*FeO	6.77
MnO	0.13(2)
NiO	0.02(2)
MgO	13.50(22)
CaO	12.19(23)
Na ₂ O	2.50(15)
K ₂ O	0.93(4)
H ₂ O	0.95(1)
F	0.30
C1	0.42(3)
-O=F	0.18
Total	98.13
Si	6.031
Al sum T	8 000
sum 1	0.000
Al	0.168
Fe ³⁺	0.327
Ti Cr	0.688
Ni	0.002
Mg	3.032
Fe ²⁺	0.781
sum C	5.000
Fe ²⁺	0.073
Mn ²⁺	0.017
Ca	1.910
sum B	2.000
Са	0.060
Na	0.731
Κ	0.179
sum A	0.970
ОН	0.865
F	0.200
0	0.935
sum W	2.000
<i>Note:</i> *FeO _{tot}	= 9.36(31) wt%; OH easured by SIMS_thic
procedure imp	lies a unique solution
for the calculat	ion of Fe^{2+}/Fe^{3+} .
Li was found level.	to be at the wt ppm

Table 6. Site populations (*apfu*) derived from X-ray and neutron site-scattering refinements for the amphibole of this work

	<i>M</i> (1)	<i>M</i> (2)	<i>M</i> (3)	Σ	EMPA
Mg	1.17	1.28	0.55	3.00	3.03
Fe	0.45	0.39	0.29	1.13	1.11
Ti	0.38	0.16	0.16	0.70	0.69
Al		0.17		0.17	0.17
Σ	2.00	2.00	1.00	5.00	5.00

Figure 1. Projection onto (100) of the structure of amphibole K(1). The site nomenclature is also shown. Structure model based on the neutron refinement of this study; displacement probability factor: 50%.



Figure 2. H-bonding configuration in the structure of K(1) amphibole deduced on the basis of the neutron structure refinement of this study. Displacement probability factor: 50%.

