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Critical Review of salinity intrusion in rivers and estuaries

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Critical Review of salinity intrusion in rivers and estuaries

Short title: Critical Review of salinity intrusion in rivers and estuaries

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Abstract

There is scientific evidence of accelerated sea level rise and saline intrusion. Some impacts, such as stratification and estuarine circulation are subtle; others are dramatic including shifts in salt-sensitive habitats and limited water availability of suitable quality for industrial and municipal uses. These results

27 have become a remarkable reality resulting in a set of integrated surface water organisation issues.
28 Tremendous population increases overwhelming many coastal areas have expanded the problem. These
29 challenges have been studied from many perspectives using various objectives and methodologies, and
30 then arriving at different findings. However, all research assured that significant rises in sea level have
31 influenced estuaries and tidally affected rivers, and these observations are expected to become rapidly
32 worse in the future. This study introduces categorises, critically investigates, and synthesises the most
33 related studies regarding accelerated sea level rise and challenges of the development associated with the
34 resources of surface water in estuaries and tidally-affected rivers. This critical review reveals that there is
35 a need for research that focuses on the development of sustainable surface water resources.

36
37 **Keywords:** Climate change; Coastal resources management; Estuary; River discharge; Tidal analysis; Water Resources
38 Management.

41 **Introduction**

43 *Background, aim, and overview*

45 Amongst all recent universal, ecological and public alterations, climate
46 variability, as estimated by many climate models (IPCC 2007), will have severe
47 influences on coastal areas. There is an extensive variety of effects containing rising sea
48 level, rain patterns, floods, and drought occurrences. These impacts may have important
49 influence on natural resources, particularly water resources (ground or surface). This is
50 mainly challenging for estuaries, where both natural and socio-economic resources of
51 great importance available and are developing quickly.

52 Estuaries, in the traditional sense, are transition zones between upland rivers and
53 coastal waters, and are critical nursery regions providing grounds for a considerable

54 number of marine animals (Zhang et al. 2012; Chen et al. 2013; Jacob et al.2013; Chen
55 et al. 2014; Liu & Liu 2014; Xu et al. 2015). The leading environmental factors
56 impacting on the survival and distribution of organisms in estuaries are salinity
57 distribution and intrusion (Attrill 2002; Chen et al. 2009; **Khadim et al. 2013**; Chua &
58 Xu 2014; Garcia et al. 2010; Hong & Shen 2012; Renaud et al. 2015). The structure of
59 salinity patterns in estuaries results from the interplay between estuarine themorphology
60 and topography, tidal elevation at the estuary mouth, saline water variation between
61 ocean and freshwater discharge. Therefore, these factor interactions determine the
62 estuarine mixing mechanisms and the saline water movement process (Shaha et al.
63 2011; Xu et al. 2015). Saline intrusion may lead to estuarine water quality decline so
64 that the corresponding water becomes inadequate for specific purposes including
65 drinking water, agricultural, and industrial uses (Renaud et al. 2015). Accordingly, the
66 simulation of saline water distribution along an estuarine system is often seen as the
67 prime concern of decision-makers in estuarine systems and coastal areas (Bhuiyan &
68 Dutta 2012; Hong & Shen 2012; Rice et al. 2012; Liu & Liu 2014).

69 Many water resources researchers are interested in salinity intrusion in estuaries,
70 and they undertook it from many perspectives, concentrating on either groundwater or
71 surface. Various tools have been applied to characterise, categorise, and analyse salinity
72 distribution in estuaries. Most of the research confirmed that sea level rise is a
73 substantial problem that should be discussed with high importance. Key references
74 related to the link between salinity intrusion and sea level rise are summarised in Table
75 1.

79

Table 1. References related to salinity intrusion and sea level rise (SLR)

Reference	Location	Finding
Kurup et al. (1998)	Swan River estuary, Australia	The freshwater river inflow is the most critical process impacting on the salt wedge site.
Liu et al. (2004)	Tanshui river system, Taiwan	The average annual salinity was 8.5 parts per thousand (ppt) and 12.8 ppt before and after the construction of the reservoir in this order.
Liu (2005)	Keelung River, Taiwan	The boundaries of saline intrusion before regulation of the channel were further upstream of the river compared to after channel regulation.
Liu et al. (2007)	Danshuei River estuarine, Taiwan	The two-layered circulation within the estuary prevails most often at the estuary mouth.
Xue et al. (2009)	Changjiang River, China	The salt-water movement resulted from a complex non-linear interaction process in relation to freshwater flow upstream, tidal currents, water mixing, wind impact, and saltwater distribution.
Jeong et al. (2010)	Modaomen Estuary, China	During neap tides, the estuary gains salt, whilst during spring tides, it loses salt.
Cai et al. (2012)	Modaomen Estuary, China	The estuary would be increasingly exposed to salt intrusion and to flooding from storm surges as a result of further deepening, which facilitates the penetration of storm surges into the system.
Bhuiyan and Dutta (2012)	Goral river network, Bangladesh	A 59-cm rise in sea level increases the salinity by 0.9 ppt at a distance of 80 km upstream of the river mouth.
Rice et al. (2012)	Chesapeake Bay, USA	With 50 cm and 100 cm sea level rises, the maximum salinity increases would be 2 and 4 ppt, respectively.
Chen et al. (2013)	Wu River estuary, Taiwan	More tidal energy will spread into the estuary after weir construction.
Mendes et al. (2013)	Douro estuary, Portugal	The amplitude of the principal lunar semi-diurnal tide constituent increased for almost the entire estuarine area
Liu and Liu (2014)	Wu River estuary, Taiwan	The overall flushing time extracted from the system for high flow without sea level rise is lower in comparison to the sea level rise.
Chen et al. (2014)	Tamsui River estuarine, Taiwan	The maximum increment of depth-averaged and tidal-averaged salinity would be 1.1, 2.4 and 3.0 psu in this order for corresponding sea level rises of 0.34, 1.05 and 1.40 m at the middle region of the estuary under the average discharge scenario.
Kuang et al. (2014)	Yangtze River Estuary	The ebb flow split ratio increases up to almost 5% under a 2-m sea level rise scenario.
Liu and Liu (2014)	Wu River estuary, Taiwan	The tidal expedition further upstream increases by 500 and 900 m under the Q_{10} and Q_{90} discharge scenarios, respectively.
Xiao et al. (2014)	Marks River Estuary, USA	A 0.85-m sea level rise caused a substantial increase in salinity near the Wakulla River with an increase of 9.2 ppt for the surface and 12.7 ppt for the bottom salinity.
Chua and Xu (2014)	No specific location (i.e.: idealised estuary model)	A sea level rise causes a strong longitudinal saline water gradient, higher longitudinal dispersion coefficients and increased saline water movement.

80

81 This research tries to specify the key outcomes of most current studies regarding
82 sea level rise and development challenges that estuaries and tidally affected rivers face
83 with especial attention to surface water. This critical review should serve as the bases
84 for specifying further research requirements. Accordingly, the current study is arranged

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85 into various sections. In the second one, recent articles concerned with the temporal and
86 spatial variations of saline water upstream movement in the estuarine systems and the
87 underlying driving mechanisms are discussed. Details associated with the responses of
88 saltwater intrusion to streamflow and tidal mixing are included. The third section
89 assessed the most important articles that investigate the impact of human interventions
90 in coastal areas, such as channel regulation and dredging, weir construction, seafloor
91 bathymetry alteration, and freshwater diversion on the inland migration of saline water.
92 The fourth section introduces the various hydrodynamic and solute transport numerical
93 models predicting the impacts of sea level rise on saline intrusion. Section **five**
94 discussed future research directions that would enhance understanding of hazards linked
95 to saline intrusion as well as help **basin** managers and decision-makers to establish
96 adaptive strategies.

97 Although this paper will investigate a sufficient number of literature sources, it
98 looks to be ultimately difficult to contain all publications in a single review paper. It is,
99 therefore, likely that some features of the theme have either been unnoticed or only
100 concisely discussed. Some of the river and estuary research saline water intrusion
101 aspects have been deliberately considered only slightly. Therefore, they merit a more
102 comprehensive specialised review. It is anticipated that these gaps could be covered by
103 following timely contributions. There is certainly target for further discussions related
104 to salinity intrusion research, probably in the broader context of future improvement of
105 total estuarine and coastal science.

110 **Factors affecting the saline intrusion**

111

112 *Freshwater flow and tidal mixing*

113

114 One requires a good understanding of river discharges and water movement

115 within an estuary to accurately predict the hydraulic transport of materials (e.g.,

116 pollutants and sediment) within streams to an estuary and then into the ocean. The two

117 hydrodynamic characteristics of tidal rivers are incorporating fluvial and bi-directional

118 flow. The former includes seasonal and transient variations of freshwater discharge,

119 while the latter is dominated by tide, and complicated by salt-freshwater inter-action as

120 well as meteorological events (Liu et al. 2001).

121 The volume of river flow is the predominant factor affecting estuarine saline

122 intrusion and plays a key role in depressing inland saltwater, particularly when river

123 flow is sufficiently high to fill the entire tidal prism completely through the phase of

124 rising tide (Liu et al. 2001; Liu et al. 2007; Graas & Savenije 2008; Gong & Shen 2011;

125 Cai et al. 2012). A growing number of studies have been conducted to investigate the

126 temporal and spatial differences of saline water movement in the estuarine system and

127 associated driving mechanisms. Among them are Kurup et al. (1998), Gibson & Najjar

128 (2000), Liu et al. (2001), Liu (2006), Liu et al. (2007), Xue et al. (2009), Becker et al.

129 (2010) and Trieu & Phong (2015). In their studies, they used different analytical and

130 hydrodynamic and solute transport models for an estuary with tributaries, in addition to

131 the main stream to investigate the change in estuarine circulation and salinity

132 distribution under different hydrological situations. They proposed that salt intrusion

133 increase within the estuary can be linked to various factors, such as upstream river

134 discharge decrease, the river mouth tidal amplitude increase, and dredging activities.

135 Their simulation results demonstrated that salt-water movement resulted from a
136 complex non-linear interaction process in relation to freshwater flow upstream, tidal
137 currents, water mixing, wind impact, and saltwater distribution. Additionally, they
138 demonstrated that river flow is the most important natural process impacting on the salt
139 wedge site and plays a major role in preventing inland movement of saline water, in
140 particular, when the flow of the tidally affected river is high. For example, they noticed
141 that the reduction of the river flow to the flow that is equal to or exceeding 75% of time
142 (Q_{75}) increases the extent of saline water intrusion. Moreover, they suggested that more
143 turbulent mixing, which in turn increases the salt intrusion, lowers vertical stratification,
144 and less residual circulation result from the spring tide.

145 The estuarine system residence time can increase with the change in river flow
146 during a scenario of rising sea level. This scenario reveals not only an alteration in
147 saline water movement, but also a rise in the time of residence. A rise in sea level is
148 likely to change the river estuary location, thereby leading to a considerable alteration
149 in both fish environment and breeding ground location. Fishes usually breed in
150 estuarine systems and develop further in saline waters. It follows that a rise in sea level
151 would shift this boundary upstream, altering the environment of fishing populations.
152 The rises in residence time as a result of sea level rise are likely to extend the
153 movement of dissolved materials, which would result in the decline of the
154 corresponding water quality.

155 Water transfer projects can effectively mitigate saline water movement. Saline
156 water intrusion and river flow discharge can be predicted faster and by a simpler
157 method through using two optimal regression equations, which provide a policy-making
158 basis allowing for a fast response to deploy corresponding preventive measures to
159 minimise risks in case when the saline water intrusion occurs.

160 Gong & Shen (2011) improved a modelling system based on nested grids to
161 simulate the saline water transport processes and the corresponding saline water
162 intrusion response to alterations in streamflow and tidal mixing in the Modaomen
163 Estuary, China. They showed that during neap tides, the estuary gains salt, whilst
164 during spring tides, it loses salt. The researchers also confirmed that the streamflow
165 pulse overturns the corresponding saline intrusion greatly.

166 All in all, the responses of saline water intrusion to streamflow and
167 corresponding tidal mixing have been researched in detail by Turrell et al. (1996),
168 Uncles & Stephens (1996), Chen (2004), Prandle (2004), Sierra et al. (2004), Brockway
169 et al. (2006), Huang (2007), and others. These studies confirm that the magnitude of
170 freshwater flow is the predominated factor impacting estuarine saline intrusion and
171 plays a key role in reducing inland saltwater, which can lead to contamination of
172 drinking water sources and other consequences.

173

174 *Manmade activities and interventions*

175

176 Many studies showed that there are various other parameters that might affect
177 the movement of saline water in coastal areas in addition to river discharge and tidal
178 level variation (Bobba 2002) such as anthropogenic interventions (e.g., river
179 regulation). For example, changes in seafloor bathymetry play a considerable role in
180 saline water movement. Therefore, to investigate the hydrodynamic characteristics and
181 salinity intrusion in the estuary in response of the bathymetry alteration, Liu et al.
182 (2004) mimiced salinity distributions in the Tanshui river, Taiwan, under various
183 bathymetric configurations and for three flow scenarios (mean flow (Q_m), the flow that
184 is equal to or exceeding 50% of time (Q_{50}) and the flow that is equal to or exceeding

185 75% of time (Q_{75}) at the upstream boundaries of three system tributaries both before
186 and after the construction of a reservoir by applying a vertical, laterally integrated, two-
187 dimensional hydrodynamic and saltwater intrusion model. Their results showed that
188 both changes have led to further intrusion of tidal flow and upstream salt-water
189 presence. Furthermore, the salinity increased throughout the year and the average
190 annual salinity was 8.5 parts per thousand (ppt) and 12.8 ppt before and after the
191 construction of the reservoir in this order. Then, Liu (2005) utilised the same model to
192 assess the change in tidal ranges, saltwater intrusion and residual circulation that result
193 from channel regulation in the Keelung River, Northern Taiwan. The researcher
194 detected that the boundaries of saline intrusion before the channel regulation were
195 further river upstream than after channel regulation, because of the impact of a deeper
196 and wider channel as well as the fundamental decline in saline water, which was
197 observed along the estuary after the channel regulation.

198 Sanders & Piasecki (2002) formulated and subsequently applied an optimisation
199 problem by a quasi-Newton method, which uses a Broyden-Fletcher-Goldfarb-Shanno
200 update to assess the gradient of the objective function with respect to the parameter
201 vector (Shanno & Phua 1980). They discovered that salinity sensitivity levels to
202 freshwater diversion from estuaries, for municipal and agricultural purposes, have both
203 intra-tidal and inter-tidal variability. Diversions at any period and at any point
204 throughout an estuary have either a long-term or a rapid impact on the salinity
205 distribution.

206 Chen et al. (2013) established and applied a three-dimensional hydrodynamic
207 water quality model to predict the possible influences of the Dadu Weir located at the
208 Wu River estuary, Taiwan, in terms of saltwater intrusion and water quality changes
209 under low flows. The model applications revealed that more tidal energy will spread

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210 into the estuary after the weir construction as a result of reduced freshwater discharges.

211 Thus, the salinity will subsequently increase.

212 The prevention of salt-water intrusion is fundamental to estuarine system
213 management. The storage system directly affects the downstream river discharge, which
214 influences salt water intrusion intensity and disturbs freshwater intake at, for example,
215 reservoirs located within the estuary boundaries. The reservoirs are normally operated
216 on a seasonal scale. Accordingly, it is important to investigate the effect of seasonal
217 regulation of streamflow by reservoirs on salt-water intrusion. Thus, a quantity
218 assessment of this critical issue is important for the active management of water
219 resources within the estuary boundaries.

220 Saline water movement enhancement and advancement has a negative influence
221 on the preservation of freshwaters. Saline intrusion becomes minimal in the estuary at a
222 time when water is released from river storage systems during dry periods.

223 Eventually, Li et al. (2014) evaluated the future availability of freshwater in the
224 Changjiang example estuary by undertaking a hydrological investigation based on the
225 examination of the available discharge, salinity, water diversion projects, and sea level
226 rise. They extrapolated salinity events into the future at intervals of ten years until the
227 year 2040. The results explained that the scarcity of estuarine freshwater scenario is
228 complicated by both the salinity intrusion through the North branch and the diversion of
229 flow from the Changjiang estuary downstream of the Datong gauging station.

230 **Human interventions in the form of tidal river management are motivated by**
231 **addressing saline water intrusion challenges in coastal regions. For example, Khadim et**
232 **al. (2013) evaluated the benefits achieved due to the implementation of integrated water**
233 **resources management strategies in some coastal regions, and assessed some technical**
234 **aspects involving a new tidal river management concept.**

235

236 **Climate change**

237

238 *Overview*

239

240 One of the most serious environmental issues facing the world is climate change.

241 It is widely accepted that global warming has the potential to affect many humans

242 dramatically and predominantly adversely as a consequence of both natural and

243 anthropogenic changes to temperature, precipitation, sea level, air quality, and other

244 climatic conditions (Mendes et al. 2013; Chen et al. 2014; Liu & Liu 2014; Scholz

245 2014).

246 Sea level rises result from three main causes: ocean thermal expansion as well as

247 mountain glaciers and ice cap melting pose a particularly disastrous threat, because they

248 have long-term impacts on coastal areas, such as increase in coastal erosion and sea

249 water intrusion (IPCC 2007; Woodworth et al. 2009). The fifth evaluation report of the

250 inter-governmental panel on climate variation explained that the average rate of global

251 mean sea level increased by 1.7 mm per year between 1901 and 2010, and it is very

252 likely that this rise will continue beyond the 21st century (IPCC 2013). For the period

253 between 2081 and 2100, their latest report expects a future rise in global sea level

254 between 0.45 m and 0.82 m for the most pessimistic scenarios (IPCC 2014).

255 A rise in sea level can lead to migration of saline water upstream in both estuaries

256 and rivers due to many factors, such as density difference, tidal range, freshwater flow,

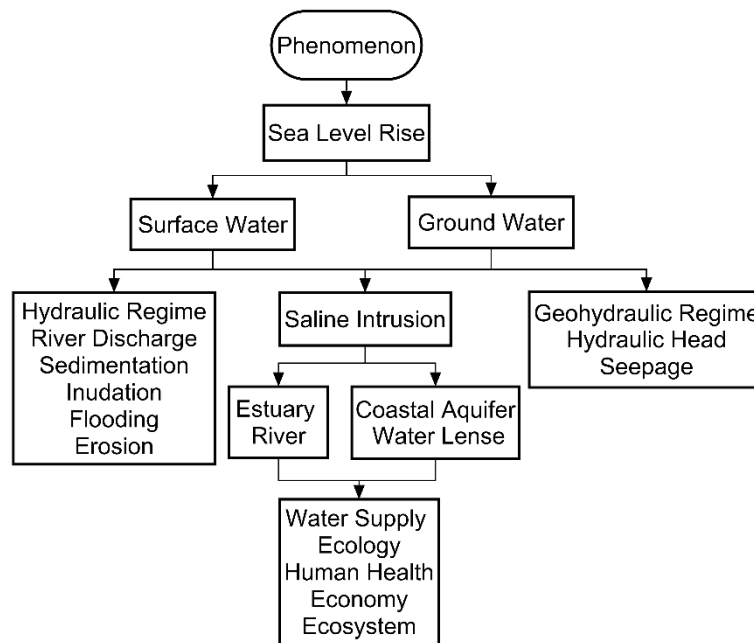
257 increase of the cross sectional area, and mixing process. Such a harmful movement may

258 deteriorate estuarine water quality, so that its water will be unsuitable for drinking,

259 agricultural and industrial purposes (Farber et al. 2005; Pathikonda et al. 2010; Bhuiyan

260 & Dutta 2012; Rice et al. 2012; Chen et al. 2013; Liu & Liu 2014). Furthermore, this
 261 could cause alterations in salt-sensitive habitats and negatively impact on the fauna and
 262 flora development. Thus, the saline water distribution alongside an estuary is a prime
 263 concern for the water management (Bhuiyan & Dutta 2012; Hong & Shen 2012; Rice et
 264 al. 2012; **Khadim et al. 2013**; Liu & Liu 2014).

265 Generally, the sea level rise impact on coastal areas water management can be
 266 divided into three parts: coastal zone and the hinterland, estuaries and rivers, and
 267 coastal groundwater flow regimes. Figure 1 reveals relevant aspects that affect the
 268 coastal areas due to sea level rises.



269
 270 **Fig. 1.** Relevant aspects that affect coastal areas due to sea level rise

271
 272 A rise in mean sea level has many consequences in coastal areas including a
 273 higher risk of low-lying coastal area flooding, sandy beach erosion, saline water
 274 intrusion and the loss of wetlands (Kirwan et al. 2008; Poulter & Halpin 2008; Gesch
 275 2009). Such impacts will lead to geomorphological, ecological, and socio-economic
 276 effects in coastal areas (Michael 2007).

277 Understanding sea level rise implications on estuaries has attracted many
278 researchers worldwide. The impact on estuaries and tidally affected rivers is of key
279 interest.

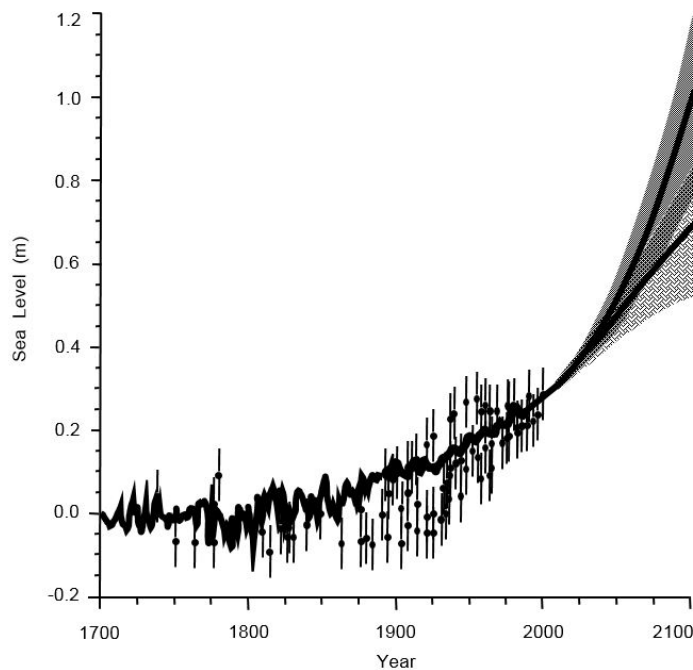
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281 *Climate change and sea level rise*

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283 One of the significant impacts of climate change is sea level rise. The fifth
284 assessment report of the IPCC (2013) explained that the average amount of global mean
285 sea level increase was 1.7 mm per annum between 1901 and 2010 (Fig. 2). Patterns and
286 trends of sea level rise can be investigated either through linear regression for long-term
287 data obtained from tidal gauges and satellite altimetry or by considering certain
288 scenarios set by the IPCC as illustrated in Fig. 3.

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291 **Fig. 2.** The global mean sea level rise average rate for the time span between 1700 and
292 2100 (after IPCC, 2013)

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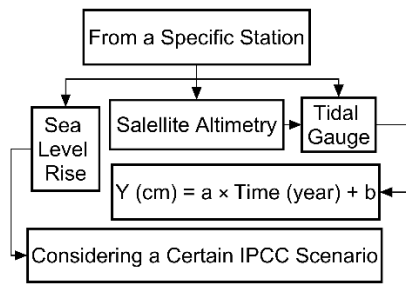


Fig. 3. Sea level rise projection. **Note:** Y represents sea level rise (cm), a and b are the regression coefficients (dimensionless), and IPCC is the Intergovernmental Panel on Climate Change

Rising sea level impacts attract many researchers' attention, particularly those focusing on the saline water change as a result of this rise in estuaries. Numerous numerical modelling has indicated that rises in sea level have a significant influence on salinity concentrations in estuaries. For example, by estimating a relative rise in sea level of 18 cm to 167 cm by 2100, Hilton et al. (2008) predicted a change of 0.4 to 12 ppt in the Chesapeake Bay salinity, USA. They applied a statistical model to simulate the monthly salinity time series from 1949 to 2006 for 23 grid cells of the bay main stream. Hilton et al. (2008) proved that the residual salinity exhibits an increase over the model run covering 57 years, indicating that there is another influential factor. The Susquehanna River flow increased the bay salinity since 1949. The potentially major future impact of sea level rise on saline water alteration would likely have disastrous impacts on estuarine ecosystems and society. Accordingly, it is essential to note that there is a critical need for further research to be conducted to achieve a better understanding of the relationship between saline water and sea level. Particularly, it is crucial to quantify alterations, both due to sedimentation and dredging, in the bathymetry of a bay, and to quantify those effects on saline intrusion by numerical modelling.

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317 **Modelling of surface water salinity intrusion**

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319 *Brief overview of available models*

320

321 Table 2 reveals that one of the most widespread codes in current centuries has
322 been the Environment Fluid Dynamics Code (EFDC). The EFDC model is a common
323 purpose modeling package for simulating three-dimensional surface flow, transport and
324 biogeochemical processes including rivers, estuaries, lakes, coastal regions, wetlands,
325 and reservoirs. It was initially established by the Virginia Institute of Marine Science as
326 authorised by US EPA. The EFDC model has three functional modules, such as
327 hydrodynamics, water quality eutrophication, and toxic sediment pollutant movement
328 and fate, entirely integrated into single software. The Mellor and Yamada Level 2.5
329 Turbulence Closure Scheme (Mellor & Yamada 1982) is implemented in the model
330 (Galperin et al. 1988). The model applies horizontal coordinates, stretched (or sigma)
331 vertical coordinates and curvilinear, orthogonal. It mimics mass and topographically
332 induced circulation in addition to tidal and wind-driven flow rates, and time-based and
333 spatial salinity distributions, temperature, and conservative/non-conservative tracers.
334 The model has a flexible grid and network structure, which is capable of connecting
335 several tributaries to the main channel through grid linkages between downstream and
336 upstream grid cells containing structures of dam.

337 Jeong et al. (2011) implemented the EFDC model to analysis the upstream
338 movement of salt water features in the downstream stretches of the Geum River basin
339 (South Korea) for four flow conditions (flood, normal, low and drought). They used
340 EFDC to investigate the influence of saline water movement when the gates of the
341 estuary barrage were fully opened. The results suggested that the maximum salt-water

342 concentration can be measured near the barrage of the estuary and the salt-water
 343 concentration declined as the distance from the barrage increased.

344

345 **Table 2.** Most widely used salinity intrusion models

Model	Basic model features	Practical salinity intrusion applications
Statistical	-	Gibson and Najjar (2000); Hilton et al. (2008); Zhang et al. (2010); Cai et al. (2012)
Linear programming	-	Sanders and Piasecki (2002)
1D ^a	-	Bhuiyan and Dutta (2012); Fleenor and Bombardelli (2013)
2D ^b	Laterally averaged Vertically averaged	Kurup et al. (1998); Hsu et al. (1999); Liu et al.(2001); Liu et al. (2004); Liu (2005)
3D ^c	MIKE 21 ^d	Kuang et al. (2014)
	MIKE 11 ^e	Dat et al. (2011)
	EFDC ^f	Jeong et al. (2010); Gong and Shen (2011); Rice et al. (2012); Hong and Shen (2012); Qiu and Zhu (2013)
	SELFE ^g	Chen et al. (2013); Liu and Liu (2014)
	Un-TRIM ^h	Liu et al. (2007)
	FVCOM ⁱ	Xue et al. (2009)
	MHIOD ^j	Mendes et al. (2013)
	POM ^k	Xiao et al. (2014)
	SUNTANS ^l	
	Chua and Xu (2014)	

^aOne dimensional.

^bTwo dimensional.

^cThree dimensional.

^dTwo-dimensional Hydrodynamic Model.

^eTwo-dimensional Hydrodynamic Model.

^fEnvironment Fluid Dynamics Code.

^gEulerian-Lagrangian Finite-Element.

^hA three-dimensional, time-dependent, baroclinic, hydrodynamic, and salinity numerical model.

ⁱFinite Volume Coastal Ocean Model.

^jA baroclinic finite volume two-dimensional numerical model.

^kPrinceton Ocean Model.

^lStanford Unstructured Non-hydrostatic Terrain-following Adaptive Navier–Stokes Simulator.

360

361 Gong & Shen (2011) improved a modelling system based on nested grids to

362 simulate the saline water transport processes and the corresponding response of salinity

363 intrusion to alterations in river flow and tidal mixing in the Modaomen Estuary, China.

364 They showed that during neap tides, the estuary gains salt, whilst during spring tides, it

1
2 365 loses salt. The researchers also confirmed that the streamflow pulse overturns the
3 366 corresponding saline water intrusion greatly.

4
5 367 The upstream saline water movement in estuaries results in limited water
6
7 368 availability of suitable quality for many purposes, such as domestic and industrial uses.
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9 369 Furthermore, under scenarios with and without the Three Georges Reservoir regulation,
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11 370 Qiu & Zhu (2013) simulated the seasonal saltwater intrusion around the Changjiang
12
13 371 Estuary (China) by applying an updated EFDC model. Throughout the dry season, they
14
15 372 found that seawater intrusion around the freshwater reservoirs considerably decreased
16
17 373 as the Three Georges Reservoir supplemented river discharge. Whereas, during autumn
18
19 374 season, this reservoir progressive the timing of salinity intrusion and marginally
20
21 375 increased its density. Accordingly, the researchers concluded that the seasonal
22
23 376 discharge regulation by the Three Georges Reservoir has both positive and negative
24
25 377 impacts on saline water intrusion and freshwater availability during the annual dry and
26
27 378 wet periods.
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34 379 To shed light on the responses of the Chesapeake Bay (USA) to sea level rise
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36 380 scenarios of the 21st century, which are set by the US Climate Change Science Program
37
38 381 (Mahoney et al. 2003), Hong & Shen (2012) utilised a three-dimensional
39
40 382 hydrodynamic-eutrophication model (HEM-3D). Based on the corresponding numerical
41
42 383 results, they pointed out that mean salt content, salinity intrusion and stratification
43
44 384 would rise during the dry season greater than in a typical year as the sea level rises.
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46 385 Both salinity concentration and stratification data reveal more likely increases in spring
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48 386 and wet years than in autumn and dry years.
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53 387 Furthermore, to explain the alteration of transport processes with rising sea levels,
54
55 388 the transport time scales were used by Hong & Shen (2012). The corresponding results
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57 389 indicated the following: Firstly, the freshwater downstream transport would be slower
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1 390 while the flow substitution would be reinforced. Secondly, owing to the greater volume
2 391 and enhanced alteration of circulation, the residence time of the bay would increase.
3
4 392 Thirdly, the vertical transport time rises and the volume of water mass for various age
5
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7 393 groups increases at different rates. Therefore, the retention time of dissolved materials
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9
10 394 would rise in the bay.

11
12 395 Additionally, by utilising EFDC, Hong & Shen (2012) demonstrated that although
13
14 396 the intensified tidal currents would increase due to vertical mixing, the strengthened
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17 397 stratification would weaken the exchange in the vertical level. As a result of
18
19 398 stratification changes, the vertical transport time would increase considerably the
20
21
22 399 impact of tidal alterations. The increased upstream transport time has less of an impact
23
24 400 on hypoxia environments in the middle and upper bay parts, because the dissolved
25
26 401 oxygen provision at the bottom of the bay is dominated by vertical exchanges. Less
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29 402 dissolved oxygen supply from the surface to the bottom region would be recorded due
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32 403 to the weakened vertical exchange. Furthermore, Rice et al. (2012) studied the sea level
33
34 404 rise impacts on tidal freshwater wetlands and on the potable water supply of two
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36 405 tributaries of Chesapeake Bay, James River and Chickahominy River using the EFDC
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39 406 model. They used sea level rise scenarios of 30, 50, and 100 cm based on the US
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41 407 Climate Change Science Program (2009; see also above) for the mid-Atlantic region
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43
44 408 concerning the 21st century (Hong & Shen 2012). For the James River, the results
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46 409 demonstrated that the salt water rises in the lower and upper reaches of the river, and
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48
49 410 are smaller than those from the middle to the upper river parts. With 50 and 100 cm sea
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51 411 level rises, the maximum saline water increase would be 2 and 4 ppt, respectively, and
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53 412 the upstream movement of the 10 ppt isohaline is much greater than the 5 and 20 ppt
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56 413 isohaline movements. Rice et al. (2012) stated that if the sea level rises by 100 cm, the
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58 414 salinity water volume rises significantly ($p < 0.05$) between 10 and 20 ppt. Whereas in
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415 the Chickahominy River, the average salinity at the abstraction point for drinking water,
416 which is located 34 km upstream of the estuary mouth, is likely to be greater than 5 ppt
417 in a dry year and nearly 3 ppt in a typical year. Furthermore, they concluded that for a
418 1m sea level rise, the James River salinity would move nearly 10 km upstream. Also,
419 Rice et al. (2012) found that during a dry year, the number of days of salinity with more
420 than 0.1 ppt would increase. For instance, 0.1 ppt would be exceeded for ≥ 100 days at a
421 small increase of 30 cm.

422 The research specifies that EFDC has been effectively applied to a wide range of
423 saline water intrusion model-based investigation issues. Similar with other codes, the
424 key drawbacks that researchers could face when applying EFDC are in computing the
425 trade-off between wanted complication that is required for interpreting the predicted
426 distribution of salinity and long running times, and the model calibration and validation
427 required efforts. Nevertheless, such modelling codes have permitted potentials for
428 simulating three-dimensional surface water and estimating the magnitudes and
429 directions of saline water intrusion under altered future circumstances.

430 431 *Surface water salinity modelling studies*

432
433 Many mathematical methods were used to evaluate and mimic the saline water
434 intrusion in the surface water. Previous research mostly concentrated on computing the
435 saline water intrusion as a consequence of sea level rise using two dimensional
436 hydrodynamic models. Examples of such studies include Kuang et al. (2014) who
437 applied a hydrodynamic model (MIKE 21; (DHI 2011)) to investigate the effects of
438 potential future sea level rises in the Yangtze River Estuary. The model simulated sea
439 level rise scenarios of 0.5, 1.0, and 2.0 m during the flood season. According to the

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440 MIKE21 model outcomes and under sea level rise conditions, they concluded the
441 following: (1) the estuarine system tidal level increases and the rate of increase declines
442 slowly upstream along the river; (2) the movements of tidal wave upstream will
443 increase resulting in the upstream progress of a tidal limit and tidal current limit; (3) the
444 flood and ebb velocities will increase; and (4) the North Branch ebb flow has the largest
445 rate of increase, with the ebb flow divided percentage increasing up to nearly 5% under
446 the 2 m sea level rise scenario.

447 However, Mendes et al. (2013) implemented a baroclinic finite volume two-
448 dimensional numerical model to explore the alterations in current velocity, amplitude
449 and phase of the main semi-diurnal and diurnal tidal constituents associated with the
450 Douro estuary as a result of the potential effect of sea level rise. The study focused on
451 three sea level projections: 0.28 m, 0.42 m and 1.00 m. The first and second scenarios
452 were addressed by Lopes et al. (2011), and the third adopted values used in several
453 studies (e.g., Sano et al. (2011); Yates et al. (2011)). The main finding of these studies
454 is that the amplitude of the principal lunar semi-diurnal tide constituent increased for
455 almost the estuarine area. This arrangement is more obvious in the intertidal region
456 close to the estuary mouth.

457 Despite the fact that two- and three-dimensional hydrodynamic and solute
458 transport models are better suited to model a complex network of tidally affected flows
459 under future sea level rises, Fleenor & Bombardelli (2013) claimed that a one-
460 dimensional model is computationally more efficient. In order to quickly perform multi-
461 year simulations for this phenomenon, they used a simplified delta network model with
462 a tidally averaged computational approach. They proved that sea level rise will increase
463 saline water throughout the Sacramento-San Joaquin Delta with time. The model is
464 capable of performing very fast simulations over a wide range of conditions, providing

465 guidance on what should be explored in depth with slower models in the future.

466 Bhuiyan & Dutta (2012) used a one-dimensional model to roughly estimate the sea

467 level rise impact in terms of salinity concentration and salt-water intrusion in the coastal

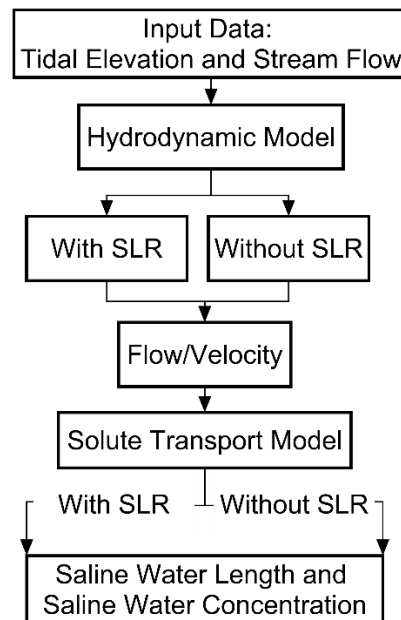
468 area of the Goral river network, South-west Bangladesh, during the dry season. The

469 overall framework of the study has been depicted in Fig. 4. They concluded that a 59

470 cm rise in sea level increases the salinity by 0.9 ppt at a distance of 80 km upstream of

471 the river mouth. This denotes a sensitivity of 0.9 ppt dividing by 0.59 m equal to 1.5 ppt

472 per meter sea level rise.



474 **Fig. 4.** Saline water intrusion modelling framework as a result of sea level rise (SLR)

475 (after Bhuiyan and Dutta, 2012)

478 Bhuiyan & Dutta (2012) stated that there is an upstream movement by nearly 21

479 km for the salinity front at 10 ppt. Moreover, the results reveal that as the flow rate in a

480 particular part is low, the saline water movement is higher in this part. Additionally,

481 Bhuiyan & Dutta (2012) claimed that despite the fact that the salinity transport

482 mechanism is well considered in computation, there is a limitation to the accuracy of

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483 this model. Although the lateral river inflow from the floodplain is considered, the
484 salinity from the floodplain is ignored. Accordingly, the floodplain lateral salinity
485 inflow needs to be inserted into the model or considered at a later stage.

486 Liu & Liu (2014) simulated the distribution of salinity and corresponding
487 transport times with respect to a sea level rise of about 38 cm in 2100 and different flow
488 rate conditions with respect to the Wu River estuary located in central Taiwan, by
489 establishing a three-dimensional hydrodynamic model entitled Eulerian-Lagrangian
490 Finite-Element (SELFE). Their experimental simulations demonstrated that the
491 intensified stratification caused a more substantial gravitation circulation increasing
492 salinity concentration by carrying a higher salinity into the estuary. The researchers
493 claimed that sea level rise increased the water surface elevation despite the fact the it
494 did not change the tide amplitude and it lengthens to the tidal expedition further
495 upstream by 500 and 900 m under the Q_{10} and Q_{90} scenarios, respectively.

496 Moreover, Liu & Liu (2014) explained that the saline water limits, which were
497 indicated by a 1 psu isohaline, are 3.0 km and 6.5 km for the current and future sea
498 level rise scenarios, respectively, under the Q_{10} condition, which is the discharge that is
499 equal to or exceeding 10% of the time. Whilst under the Q_{90} condition, such limits were
500 5.50 km and 8.25 km without and with sea level rise, respectively. The model findings
501 also indicated that the overall flushing time, which offers an assessment of the time
502 over which pollutants are released into the estuary, extracted from the system for high
503 flow without sea level rise is lower in comparison to the sea level rise, but the
504 corresponding flushing time is higher without sea level rise for low flow, if compared to
505 the rise in sea level. The results showed that a sea level rise level indicates not only a
506 variation in salinity intrusion, but also a residence time proliferation, extending the

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2 507 estuary dissolved substance transport, and subsequently leading to a deterioration of the
3 508 corresponding water quality.

4
5 509 The effects of three sea level rise scenarios (0.34, 1.05, and 1.40 m for the year
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7 510 2100) impacting on salinity, residence time and the water age of dissolved materials in
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9 511 the Tamsui River estuarine and the nearby coastal sea of northern Taiwan were assessed
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11 512 by Chen et al. (2014). For this partially mixed estuary, a three-dimensional semi-
12
13 513 implicit Eulerian-Lagrangian finite-element numerical (SELFE) model has been
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15 514 utilised. The prime finding is that the water age will rise in response to the sea level
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17 515 rise, because the dissolved materials concentration has a longer transport time from the
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19 516 upstream to the downstream locations. This is due to the increase of water volume with
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21 517 sea level rise. The residence time of the total system would also increase by nearly 17%
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23 518 under low flow conditions. In addition, the results revealed that with sea level rise, there
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25 519 will be an increase in the mean salt content and salt intrusion length. The salinity limit
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27 520 advances further toward the reaches in the upstream. The results also indicate that the
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29 521 extreme increment of depth-averaged and tidal-averaged salinity would be 1.1, 2.4, and
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31 522 3.0 psu, respectively for the sea level rises of 0.34, 1.05, and 1.40 m at the middle of the
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33 523 estuary under the average discharge scenario. Finally, the regression between length of
34
35 524 salinity intrusion and upstream freshwater flow are determined corresponding to various
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37 525 sea level rise developments.

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39 526 The impact of sea level rise on saline water movement in the Saint Mark River
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41 527 estuary has been studied by Xiao et al. (2014). They utilised a three-dimensional
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43 528 hydrodynamic model. They explained that a 0.85-m sea level rise caused a substantial
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45 529 increase in salinity near the Wakulla River with an increase of 9.2 ppt for the surface
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47 530 and 12.7 ppt for bottom salinity. Understanding the potential impacts of sea level rise
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49 531 on saline water intrusion is critical for the environmental and water resources
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5 532 management community to develop climate-adapted management plans and mitigation
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7 533 measures to protect ecosystems.

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10 534 To study the sea level rise and river flow impacts on estuarine circulation, Chua &
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12 535 Xu (2014) used an idealised estuary model. Firstly, they pointed out that a sea level rise
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14 536 causes a strong longitudinal saline water gradient, indicating an increase in the strength
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16 537 of the gravitational circulation, higher longitudinal dispersion coefficients and increased
17
18 538 saline water movement. Secondly, under low-flow conditions, the impacts of sea level
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20 539 rise on salinity intrusion are largest since the sea level rise has a greater effect owing to
21
22 540 weaker vertical stratification. Thirdly, a high flow leads to an increase of the
23
24 541 gravitational circulation, resulting in large vertical stratification, which causes non-
25
26 542 linear feedback between vertical mixing and stratification.

27 543 The low-lying Hau River is directly affected by sea level rise owing to global
28
29 544 warming. Therefore, to evaluate these impacts on the hydraulic regime and saline water
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31 545 intrusion concerning the Hau River, Doung et al. (2015) developed a two-dimensional
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33 546 hydrodynamic and solute transport model using MIKE-21 (DHI 2011). Yang et al.
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35 547 (2015) combined a watershed model with a hydrodynamic model to predict the impacts
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37 548 of climate change as well as land use and land cover change for both increase in
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39 549 freshwater flow, delivered from snowpack and precipitation, and sea level rise in the
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41 550 Snohomish estuary, north-west Washington State, USA. Huang et al. (2015) studied the
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43 551 potential of future sea level rise on saline water distribution and oyster growth in
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45 552 Apalachicola Bay using wind, tide and flow data for the period between 10 June and 9
46
47 553 July 2005. They examined the sea level rise impacts (0.31 m, 0.50 m, and 1.00 m)
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49 554 coupled with a wide range of flow conditions, such as minimum, average and maximum
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51 555 monthly flow based on the data from 1977 to 2013. Salinity movement under low flow
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2 556 is much higher than that under average flow, while salinity under high flow is much
3 557 lower than that under mean condition.

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5 558 Long-term salinity records for the time span between 1950 and 2015 were
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7 559 gathered by the Haskin Shellfish Research Laboratory and the U.S. Geological Survey
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9 560 and, and coupled with non-parametric statistical models by Ross et al. (2015) to assess
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11 561 the climate change impact on salinity distribution in the Delaware estuarine. The model
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13 562 results indicated that while insignificant trends were found at points that are usually
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15 563 upstream of the salt front, various points along the estuary show significantly increasing
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17 564 trends in salinity. In addition, the models indicated a positive correlation between sea
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19 565 level rise and enhancing residual salinity. Additionally, the findings confirmed that
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21 566 wind stress plays a role in driving salinity distributions, consistent with its impact on
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23 567 vertical mixing and Ekman transport (Banas et al. 2004). The latter is part of the Ekman
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25 568 motion theory (Jenkins & Bye 2006). The findings show that future rises in sea level
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27 569 will increase salinity, irrespectively of any alteration in the streamflow.
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34 570 An increase in knowledge of the potential impact sea level rise on coastal areas
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36 571 and estuaries is important. A sea level rise has a direct effect on salinity intrusion,
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38 572 which in turn can threaten freshwater habitats and drinking water supplies in these
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40 573 areas, which would have disastrous impacts on estuarine ecosystems and society.
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42 574 Furthermore, comprehensive scientific research presenting the amount of climate
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44 575 variation and sea level rise bearing on saline water intrusion compared to other features,
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46 576 such as water abstraction do not exist (Jackson et al. 2013; Mabrouk et al. 2013; Pervez
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48 577 & Henebry 2015). This indicates the need for future research that will assess the sea
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50 578 level rise impact versus extensive abstraction as another factor for salinity intrusion
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52 579 causes.
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581 **Avenues for future research**

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583 This research concludes with numerous ideas about the potential future directions that
584 could offer valuable inputs for the surface water resources sustainable management
585 from coastal areas. This review is part of an ongoing process to improve understanding
586 of the potential impacts of salinity intrusion in rivers and estuaries, as well as to identify
587 the key scientific questions that need to be addressed in the future. Although there is
588 consensus about the human interventions and sea level rise impacts on salinity
589 intrusion, there is less certainty about the integration between these likely impacts on
590 saline intrusion, especially due to regional changes in sea level rise. However, it is not
591 too early for basin managers and decision-makers to consider long-term adaptation
592 strategies, taking into consideration the following:

593 They should firstly investigate the positive and negative effects of river regulation
594 on salinity intrusion and freshwater resources during dry and wet periods. Water
595 impoundment in a man-made reservoir is the main cause for reduction in streamflow
596 and strengthened saline water intrusion throughout the wet season. In comparison,
597 water yielded during dry periods dilutes the salinity around the mouth of the river. This
598 is why it is vital to estimate the effects of reservoir regulation and river discharge on the
599 concentration and dispersal of saltwater intrusion in combination with the impacts that
600 might result from a future sea level rise. The dissimilarity between residual water
601 transport before and after river regulation during spring-neap tide will change.
602 Therefore, the focus should be on the spatial and temporal variations of saltwater
603 intrusion, and the impact of seasonal change of river flow on water intakes at the tidally
604 affected reaches.

605 In the absence of adaptation, saline water intrusion will gradually lead to the
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2 606 decrease in estuarine water quality, so that the corresponding water becomes
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4 607 inappropriate for many uses, such as drinking water, industrial, and agricultural
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7 608 purposes, subsequently causing an increase in poverty, potential population
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10 609 displacement and health issues. However, specific adaptation measures for a
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12 610 geographical region cannot systematically be exported to another coastal environment
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14 611 due to different geophysical, economic, social, cultural, and political regulations.
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17 612 Nevertheless, the general principles for adapting agro-ecosystems to more saline
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19 613 environments, developing crop varieties that are more tolerant to salinity, limiting
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21 614 salinity intrusion or optimising flows and sediment transport upstream through
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24 615 infrastructure development and management, restoring previously degraded coastal
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26 616 ecosystems, and institutional and governance issues, are relevant to most deltas.

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29 617 Given that the key drivers for more declining effects on estuaries and tidally
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31 618 affected rivers are identified to be sea level rise and increased development-driven
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34 619 surface water abstraction, more investigation with the developed numerical salinity
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36 620 models should be carried out on quantifying variations in surface water and water
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39 621 budget. Scenarios of current climate variation could be applied to formulate potential
40
41 622 future sea level and hydrological conditions; while a regional development plans could
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43 623 offer information for future levels and spatial distribution estimation of surface water
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46 624 abstractions.

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49 625 Exceptional importance should be put on combinations of sea level rise and
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51 626 surface water abstractions and extreme conditions. Likely future variations in the tidally
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53 627 affected rivers at the inflow to the estuary should be inspected as these would be
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56 628 reflected in different boundary conditions of the salinity intrusion models (Mabrouk et

629 al. 2013). The outcomes of the model-based analysis could serve as main objectives for
630 the introduction of probable future moderation and adaptation policies.

631 The probable effects of moderation and adaptation measures on the basis of the
632 salinity intrusion model-generated results could then be quantified and various
633 measures could subsequently be assessed. Strategies for surface water conditions should
634 be evaluated with and without mitigation scenarios, and the most appropriate scenarios
635 could then be suggested for application.

636 Thirdly, managers should search for suitable ways to reduce the effects of rising
637 sea level on saline water intrusion. This might include the construction of weirs or a
638 series of weirs to prevent saline intrusion in rivers and estuaries.

639 Additionally, it would be beneficial to apply computer models to investigate
640 salinity hazards due to an upstream shift of the brackish water zone, resulting from both
641 future rise in sea level and the increase in freshwater diversion from estuarine systems
642 for agricultural, municipal, and industrial uses.

643 Furthermore, while it is commonly accepted that a sea level rise will change the
644 spatial and temporal distribution of salinity in estuaries, it is not well-known how the
645 estuarine ecology and the abstraction points of drinking water supply will be altered and
646 affected. Further research is recommended on quantifying the association between sea
647 level rise and salinity. Particularly, it will be important to improve the quantification of
648 alterations in the bay bathymetry, due to both sedimentation and subsequent dredging.
649 Moreover, the quantification of those effects on salinity through modelling and
650 subsequent application to other estuaries of the statistical and hydrodynamic modelling
651 techniques (presented in this critical review article) is also advantageous.

652 **It** is important to note that there is a critical need for further work to be
653 undertaken to achieve an improved understanding of the link between sea level and

654 saline water. Particularly, the quantification of changes, both due to sedimentation and
655 dredging, in a bay bathymetry, and corresponding effects on saline intrusion by
656 modelling is crucial. There is also a need for a sensitivity analysis to identify the
657 conditions under which the key alterations in the saline water toe are experienced for
658 slight variations in the main hydrogeological variables, and to quantify the location of
659 the toe relative to any recharge dependent water table level.

660 A decline of fluvial flow in the tidally affected rivers would lead to more intense
661 saline intrusion. Investigations to assess the respond of coastal environments to sea
662 level rise are necessary to better understand water availability for future climate
663 scenarios.

664 Saline sub-surface water and surface water interactions are a common in the
665 coastal zone of many areas throughout the world, but the hydrological processes and
666 physiographic reasons involved are not totally understood. Further research work is
667 critical to identify such processes and factors that control the spatiotemporal variability
668 and dynamics of this interaction due to climate change and sea level rise, and
669 consequently the involvement of saline sub-surface water to surface water salinity.

670 Eventually, owing to regional circulation patterns or to vertical land movements,
671 which can be of a similar order (mm/year) of magnitude as sea level rises, the mean of
672 sea level rise can differ dramatically from the universal mean. The unequal alteration of
673 the sea level all over the world is apparent from tidal gauge and satellite altimetry
674 recordings. It follows that it becomes vital to take into account the findings from
675 regional investigations. Although considerable research has been devoted to climate
676 change effects on many estuaries universally, little consideration has been given to
677 these effects on local estuaries.

678

679 **Conclusions and recommendations**

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681 According to this critical review, sea level rise and its impacts on estuaries and
682 tidally affected rivers were the subject of many comprehensive studies going back as
683 early as 1845. Most research studies concentrated on estimating the effect of sea level
684 rise on salinity intrusion regarding surface water. However, more research is needed as
685 the results from these studies are far from conclusive.

686 Research of sea level rise due to climate change has commonly concentrated on
687 computing effects on coastal areas and surface water. Impacts of sea level rise on
688 groundwater in terms of increased saline water intrusion have been evidently
689 recognised, but quantification of these effects is currently missing. An incorporated
690 groundwater model of the coastal area aquifer that includes freshwater-saltwater
691 interactions could serve as a tool for quantification and characterisation of these effects.

692 Increased and fundamentally uncontrolled groundwater and surface water
693 abstractions are a possibly serious threat to the salinisation of the coastal aquifers and
694 surface waters. Historical trends demonstrate increases of surface and ground water
695 abstractions since 1845 predominantly modelling studies reported in literature simulate
696 coastal area deterioration as well as salinitation of the tidally affected rivers and
697 aquifers. At the same time, however, the majority of reported modelling studies are of
698 local nature, implemented in specific regions to analyse the problems of a particular
699 zone and interpret the results in terms of impacts caused by abstractions.

700 Past modelling effort was primarily carried out using two-dimensional vertical or
701 horizontal models. A significant restriction of vertical models is that the representative
702 cross sections should be selected carefully. The horizontal models provide spatial

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2 703 results in horizontal dimensions, but they give less accurate locations and shapes of the
3 704 transition zone between fresh and saltwater in the vertical dimension.

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5 705 This condition specifies that future studies should concentrate on the
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7 706 improvement of three-dimensional models, but owing to their intricacy, data
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9 707 requirements and long computational times, such models are hardly developed for
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11 708 seawater intrusion problems. Yet, such models are evidently required: Firstly, they can
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13 709 be used for hypothesis testing and better understanding of the overall system behaviour.
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15 710 Secondly, these models can be used for an integrated assessment of all potential threats
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17 711 salinisation of the area. Finally, once completely advanced, such models can become
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19 712 fundamental works of future preparation platforms and decision systems for assessment
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21 713 of many adaptation and moderation measures.
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30
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37 719 Limoges) for their kind help and support.
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44 721 **Notation**

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48 723 *The following symbols are used in this paper:*

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51 724 EFDC = Environmental Fluid Dynamics Code;

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53 725 Flushing time = The time needed to replace the estuary freshwater volume at the rate of
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55 726 the net flow;

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58 727 HEM-3D = Hydrodynamic-Eutrophication Model;

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5 730 Hypoxia = Reduced oxygen content of air or a body of water detrimental to aerobic
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7 731 organisms;
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9
10 732 MIKE21 = Two-dimensional Hydrodynamic Model;
11
12 733 Neap tide = A tide in which the variation between high and low tide is the least;
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14 734 ppt = Parts per thousand;
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16
17 735 psu = Practical salinity unit;
18
19 736 Q_{10} = The discharge that is equal to or exceeding 10% of time;
20
21
22 737 Q_{75} = The discharge that is equal to or exceeding 75% of time;
23
24 738 Q_{90} = The discharge that is equal to or exceeding 90% of time;
25
26
27 739 Q_m = Mean river flow;
28
29 740 SELF = Semi-implicit Euler-Lagrange Finite-element;
30
31
32 741 Spring tide = A tide that occurs when the difference between high and low is;
33
34 742 Tidal amplitude = The elevation of tidal high water above mean sea level;
35
36
37 743 Un-TRIM = A three-dimensional, time-dependent, baroclinic, hydrodynamic, and
38
39 744 salinity numerical model; and
40
41 745 X_{L0} = The calibrated intrusion length.
42
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46 747 **References**

47
48
49 748
50
51 749 Attrill M.J. 2002 A testable linear model for diversity trends in estuaries. *Journal of Animal Ecology*,
52
53 750 **71**(2), 262–269.
54
55 751 Banas N., Hickey B., MacCready P. & Newton J. 2004 Dynamics of Willapa Bay, Washington: a highly
56
57 752 unsteady, partially mixed estuary. *Journal of Physical Oceanography*, **34**(11), 413–427.
58
59
60
61
62
63
64
65

- 753 Bhuiyan M.J.A.N. & Dutta D. 2012 Assessing impacts of sea level rise on river salinity in the Gorai
 1 River network, Bangladesh. *Estuarine, Coastal and Shelf Science*, **96**, 219–227.
 2
 3
 4 755 <http://dx.doi.org/10.1016/j.ecss.2011.005>.
 5
 6 756 Becker M.L., Jr R.A.L. & Mallin M.A. 2010 Hydrodynamic behavior of the Cape Fear River and
 7
 8 757 estuarine system: A synthesis and observational investigation of discharge-salinity intrusion
 9
 10 758 relationships. *Estuarine, Coastal and Shelf Science*, **88**, 407–418.
 11
 12 759 <http://dx.doi.org/10.1016/j.ecss.2010.04.022>.
 13
 14 760 Bobba A.G. 2002 Numerical modelling of salt-water intrusion due to human activities and sea-level
 15
 16 761 change in the Godavari Delta, India. *Hydrological Sciences Journal*, **47**(S1), S67–S80.
 17
 18 762 <http://dx.doi:10.1080/02626660209493023>.
 19
 20 763 Brockway R., Bowers D., Hogueane A., Dove V. & Vassele V. 2006 A note on salt intrusion in funnel-
 21
 22 764 shaped estuaries: application to the Incomati estuary. Mozambique. *Estuarine, Coastal and Shelf*
 23
 24 765 *Science*, **66**(1–2), 1–5. <http://dx.doi:10.1016/j.ecss.2005.07.014>.
 25
 26 766 Cai H., Savenije H.H.G., Yang Q., Ou S. & Lei Y. 2012 Influence of river discharge and dredging on
 27
 28 767 tidal wave propagation: Modaomen estuary case. *Journal of Hydrologic Engineering*, **138**(10),
 29
 30 768 885–896. [http://dx.doi:10.1061/\(ASCE\)HY.1943-7900.0000594](http://dx.doi:10.1061/(ASCE)HY.1943-7900.0000594).
 31
 32 769 Chen S.S., Fang L.G., Li H.L., Zhang L.X. & Huang W.R. 2009 Remote sensing of turbidity in seawater
 33
 34 770 intrusion reaches of Pearl River Estuary - A case study in Modaomen water way, China.
 35
 36 771 *Estuarine, Coastal and Shelf Science*, **82**(1), 119–127.
 37
 38 772 Chen W.B., Liu W.C. & Hsu M.H. 2014 Modeling assessment of a saltwater intrusion and a transport
 39
 40 773 time scale response to sea-level rise in a tidal estuary. *Environmental Fluid Mechanics*, **15**(3),
 41
 42 774 491–514. <http://dx.doi:10.1007/s10652-014-9367-y>.
 43
 44 775 Chen W.B., Liu W.C. & Huang L.T. 2013 The influence of weir construction on salt water intrusion and
 45
 46 776 water quality in a tidal estuary-assessment with modeling study. *Environmental Monitoring*
 47
 48 777 *Assessment*, **185**(10), 8169–8184. <http://dx.doi:10.1007/s10661-013-3165-8>.
 49
 50 778 Chen X.J. 2004 Modeling hydrodynamics and salt transport in the Alafia River estuary, Florida during
 51
 52 779 May 1999 – December 2001. *Estuarine, Coastal and Shelf Science*, **61**(3), 477–490.
 53
 54 780 <http://dx.doi:10.1016/j.ecss.2004.06.012>.
 55
 56
 57
 58
 59
 60
 61
 62
 63
 64
 65

- 781 Chua V.P. & Xu M. 2014 Impacts of sea-level rise on estuarine circulation: an idealized estuary and San
1
2 782 Francisco Bay. *Journal of Marine Systems*, **139**, 58–67.
3
4 783 <http://dx.doi:10.1016/j.jmarsys.2014.05.012>.
5
6 784 DHI (Danish Hydraulic Institute) 2011 *MIKE 2I and MIKE 3 Flow model*. Flow model hydrodynamic
7
8 785 and transport module scientific documentation. Horsholm, Denmark, 2–4.
9
10 786 Farber E., Vengosh A., Gavrieli I., Marie A., Bullen T.D., Mayer B., Holtzman R., Segal M. & Shavit U.
11
12 787 2005 Management scenarios for the Jordan River salinity crisis. *Applied Geochemistry*, **20**(11),
13
14 788 2138–2153. <http://dx.doi:10.1016/j.apgeochem.2005.07.007>.
15
16 789 Fleenor W.E. & Bombardelli F.A. 2013 Simplified 1-D hydrodynamic and salinity transport modelling of
17
18 790 the Sacramento-San Joaquin delta: sea level rise and water diversion effects. San Francisco
19
20 791 *Estuary and Watershed Science*, **11**(4), 1–22.
21
22 792 Galperin B., Kantha L.H., Hassid S. & Rosati A. 1988 Aquasi-equilibrium turbulent energy model for
23
24 793 geophysical flows. *Journal of Atmospheric Sciences*, **45**, 55–62.
25
26 794 Garcia A., Juanes J.A., Alvarez C., Revilla J.A. & Medina R. 2010 Assessment of the response of a
27
28 795 shallow macrotidal estuary to changes in hydrological and wastewater inputs through numerical
29
30 796 modelling. *Ecological Modelling*, **221**(8), 1194–1208.
31
32 797 Gesch D.B. 2009 Analysis of LIDAR elevation data for improved identification and delineation of lands
33
34 798 vulnerable to sea-level rise. *Journal of Coastal Research*, SI 53, 50–59.
35
36 799 Gibson J.R. & Najjar R.G. 2000 The response of Chesapeake Bay salinity to climate-induced changes in
37
38 800 streamflow. *American Society of Limnology and Oceanography*, **45**(8), 1764–1772.
39
40 801 Gong W. & Shen J. 2011 The response of salt intrusion to changes in river discharge and tidal mixing
41
42 802 during the dry season in the Modaomen Estuary, China. *Continental Shelf Research*, **31**(7–8),
43
44 803 769–788. <http://dx.doi:10.1016/j.csr.2011.01.011>.
45
46 804 Graas S. & Savenije H.H.G. 2008 Salt intrusion in the Pungue estuary, Mozambique: effect of sand banks
47
48 805 as natural temporary salt intrusion barrier. *Hydrology and Earth System Science*, **5**, 2523–2542.
49
50 806 Hilton T.W., Najjar R.G., Zhong L., Li M. 2008 Is there a signal of sea-level rise in Chesapeake Bay
51
52 807 salinity? *Journal Geophysical Research*, **113**(C09002), 1–12.
53
54 808 <http://dx.doi:10.1029/2007JC004247>.
55
56
57
58
59
60
61
62
63
64
65

- 809 Hong B. & Shen J. 2012 Responses of estuarine salinity and transport processes to potential future sea-
1
2 810 level rise in the Chesapeake Bay. *Estuarine, Coastal and Shelf Science*, **104–105**, 33–45.
3
4 811 <http://dx.doi.org/10.1016/j.ecss.2012.03.014>.
5
6 812 Huang W. 2007 Hydrodynamic modeling of flushing time in a small estuary of North Bay, Florida.
7
8 813 *Journal Estuarine and Coastal Shelf Science*, **74**(4), 722–731.
9
10 814 Huang W., Hangen S., Bacopoulos P. & Wang D. 2015 Hydrodynamic modeling and analysis of sea-
11
12 815 level rise impacts on salinity for oyster growth in Apalachicola Bay, Florida. *Estuarine, Coastal*
13
14 816 *and Shelf Science*, **156**, 7–18. <http://dx.doi.org/10.1016/j.ecss.2014.11.008>.
15
16 817 IPCC 2007 The Physical Science Basis, in: *Contribution of Working Group I to the Fourth Assessment*
17
18 818 *Report of the Intergovernmental Panel on Climate Change*, Solomon S, Qin D., Manning M., Z.,
19
20 819 Marquis M., Averyt K.B., Tignor M. & Miller H.L. (Eds.), Cambridge University Press,
21
22 820 Cambridge, UK and New York, NY, USA, 1-103.
23
24 821 IPCC 2013 The physical science aspects basis, in: *Contribution of Working Group I to the Fifth*
25
26 822 *Assessment Report of the Intergovernmental Panel on Climate Change*, Stocker T. A., Qin D.,
27
28 823 Plattner G.-K., Tignor M.M.B., Allen S.K., Boschung J., Nauels A., Xia Y., Bex V., Midgley P.M.
29
30 824 (Eds.), Cambridge University Press, Cambridge, New York, NY, 1-1535.
31
32 825 IPCC 2014 Assessment of physical scientific aspects of climate change, in: *Contribution of Working*
33
34 826 *Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*,
35
36 827 Edmonds, A., Holy, N. (Eds.), Cambridge University Press, Cambridge, New York, NY, USA.
37
38 828 Jackson C.R., Mackay J.D. & Bloomfield J.P. 2013 Changes in groundwater levels in the UK over the
39
40 829 21st century: an assessment of evidence of impacts from climate change. Living with
41
42 830 Environmental Change (A climate change report card for water). <http://nora.nerc.ac.uk/503393>
43
44 831 (accessed 18 October 2016).
45
46 832 Jacob B., Revichandran C. & Kumar N.K.R. 2013 Salt intrusion study in Cochin estuary using empirical
47
48 833 model. *Indian Journal of Geo-Marine Sciences*, **42**(3), 304–313.
49
50 834 Jenkins A.D. & Bye J.A.T. 2006 Some aspects of the work of V.W. Ekman. *Polar Record*, **42**(220), 15–
51
52 835 22. <http://dx.doi.org/10.1017/S0032247405004845>.
53
54 836 Jeong S., Yeon K., Hur Y. & Oh K. 2011 Salinity intrusion characteristics analysis using EFDC model in
55
56 837 the downstream of Geum River. *Journal of Environmental Sciences*, **22**(6), 934–939.
57
58
59
60
61
62
63
64
65

- 838 Khadim F.K., Kar K.K., Halder P.K., Rahman M.A. & Morshed A.M. 2013 Integrated water resources
1 management (IWRM) impacts in south west coastal zone of Bangladesh and fact-finding on tidal
2 839 river management (TRM). *Journal of Water Resource and Protection*, **5**(10), 953–961.
3
4 840
5
6 841 Kirwan M.L., Murray A.B. & Boyd W.S. 2008 Temporary vegetation disturbance as an explanation for
7 permanent loss of tidal wetlands. *Geophysical Research Letters*, **35**, L05403.
8 842
9 05410.01029/02007GL032681.
10 843
11
12 844 Kotera A., Nguyen D.V., Sakamoto T., Iizumi T. & Yokozawa M. 2014 A modeling approach for
13 assessing rice cropping cycle affected by flooding, salinity intrusion, and monsoon rains in the
14 845 Mekong Delta, Vietnam. *Paddy Water Environment*, **12**, 343–354.
15 846
16 847 <http://dx.doi.org/10.1007/s10333-013-0386-y>.
17
18 848 Kuang C., Chen W., Gu J., Zhu D.Z., He L. & Huang H.C. 2014 Numerical assessment of the impacts of
19 potential future sea-level rise on hydrodynamics of the Yangtze River Estuary, China. *Journal of*
20 849 *Coastal Research*, **30**(3), 586–597. <http://dx.doi.org/10.2112/JCOASTRES-D-13-00149.1>.
21 850
22
23
24 851 Kurup G.R., Hamilton D.P. & Patterson J.C. 1998 Modelling the effect of seasonal flow variations on the
25 position of salt wedge in a microtidal estuary. *Estuarine, Coastal and Shelf Science*, **47**(2), 191–
26 852 208. <http://dx.doi:10.1006/ecss.1998.0346>.
27
28
29
30 853
31
32 854 Li M., Chen Z., Finlayson B., Wei T., Chen J., Wu X., Xu H., Webber M., Barnett J. & Wang M. 2014
33 Water diversion and sea-level rise: potential threats to freshwater supplies in the Changjiang river
34 855 estuary. *Estuarine, Coastal and Shelf Science*, **156**, 52–60.
35 856
36 857 <http://dx.doi:10.1016/j.ecss.2014.07.007>.
37
38
39
40 858 Liu W.C., Hsu M.H., Wu C.R., Kuo A.Y. & Kuo J.T. 2001 The influence of river discharge on salinity
41 intrusion in the Tanshui Estuary, Taiwan. *Journal of Coastal Research*, **17**(3), 544–552.
42 859
43
44 860 Liu W.C. 2005 Effects of channel regulation on salt intrusion and residual circulation of Keelung River.
45 *Hydrological Process*, **19**(20), 4039–4054. <http://dx.doi:10.1002/hyp.5870>.
46 861
47
48 862 Liu W.C. 2006 Modelling circulation and vertical mixing in estuaries. Proceeding-institution of civil
49 engineers maritime engineering. *Proceedings of the Institution of Civil Engineers*, **159**(2), 67–76.
50 863
51 864 <http://dx.doi:10.1680/maen.2006.159.2.67>.
52
53
54 865 Liu W.C., Chen W.B., Cheng R.T., Hsu M.H. & Kuo A.Y. 2007 Modeling the influence of river
55 discharge on salt intrusion and residual circulation in Danshuei River estuary, Taiwan. *Continental*
56 866 *Shelf Research*, **27**(7), 900–921. <http://dx.doi:10.1016/j.csr.2006.12.005>.
57
58
59
60
61
62
63
64
65

- 868 Liu W.C., Hsu M.H., Wu C.R., Wang C.F. & Kuo A.Y. 2004 Modeling salt water intrusion in Tanshui
1 River Estuarine System — Case-study contrasting now and then. *Journal of Hydrologic*
2 *Engineering*, **130**(9), 849–859. [http://dx.doi:10.1061/\(ASCE\)0733-9429\(2004\)130:9\(849\)](http://dx.doi:10.1061/(ASCE)0733-9429(2004)130:9(849)).
3
4 870
5
6 871 Liu W.C. & Liu H.M. 2014 Assessing the impacts of sea level rise on salinity intrusion and transport time
7 scales in a tidal estuary, Taiwan. *Water*, **6**(2), 324–344. <http://dx.doi:10.3390/w6020324>.
8
9
10 873 Lopes C.L., Silva P.A., Dias J.M., Rocha A., Picado A., Plecha S. & Fortunato A.B. 2011 Local sea level
11 change scenarios for the end of the 21st century and potential physical impacts in the lower Ria de
12 Aveiro (Portugal). *Continental Shelf Research*, **31**, 1515–1526.
13
14 875
15
16 876 Mabrouk M.B., Jonoski A., Solomatine D. & Uhlenbrook S. 2013 A review of seawater intrusion in the
17 Nile Delta groundwater system – the basis for assessing impacts due to climate changes and water
18 877 resources development. *Hydrology and Earth System Sciences*, **10**, 10873–10911.
19
20 878
21
22 879 <http://dx.doi:10.5194/hessd-10-10873-2013>.
23
24 880 Mellor G.L. & Yamada T. 1982 Development of a turbulence closure model for geophysical fluid
25 problems. *Reviews of Geophysics*, **20**, 851–875.
26
27 881
28 882 Mendes R., Vaz N. & Dias M.J. 2013 Potential impacts of the mean sea level on the hydrodynamics of
29 the Douro river estuary. *Journal of Coastal Research*, **65**(special issue), 1951–1956.
30
31 883
32 884 <http://dx.doi:10.2112/SI65-330.1>.
33
34 885 Michael J.A. 2007 Episodic flooding and the cost of sea-level rise. *Ecological Economics*, **63**, 149–159.
35
36 886 Pathikonda S., Meerow A., Zhenxiang H. & Mopper S. 2010 Salinity tolerance and genetic variability in
37 freshwater and brackish *Iris hexagona* colonies. *American Journal of Botany*, **97**(9), 1438–1443.
38
39 887
40 888 <http://dx.doi:10.3732/ajb.0900356>.
41
42 889 Pervez M.S. & Henebry G.M. 2015 Assessing the impacts of climate and land use and land cover change
43 on the fresh water availability in the Brahmaputra River basin. *Journal of Hydrology: Regional*
44 *Studies*, **3**, 385–411.
45
46 890
47
48 891 Poulter B. & Halpin P.N. 2008 Raster modelling of coastal flooding from sea-level rise. *International*
49 *Journal of Geographical Information Science*, **22**, 167–182.
50
51 892
52
53 893 Prandle D. 2004 Saline intrusion in partially mixed estuaries. *Estuarine, Coastal and Shelf Science*, **59**(3),
54 385–397. <http://dx.doi:10.1016/j.ecss.2003.10.001>.
55
56
57
58
59
60
61
62
63
64
65

- 896 Qiu C. & Zhu J.R. 2013 Influence of seasonal runoff regulation by the Three Gorges Reservoir on
1 saltwater intrusion in the Changjiang River Estuary. *Conventional Shelf Research*, **71**, 16–26.
2
3
4 898 <http://dx.doi.org/10.1016/j.csr.2013.09.024>.
- 5
6 899 Renaud F.G., Le T.T.H., Lindener C., Guong V.T. & Sebesvari Z. 2015 Resilience and shifts in agro-
7 ecosystems facing increasing sea-level rise and salinity intrusion in Ben Tre Province, Mekong
8 900 Delta. *Climate Change*, **133**, 69–84. <http://dx.doi.org/10.1007/s10584-014-1113-4>.
9 901
- 10
11 902 Rice K.C., Hong B. & Shen J. 2012 Assessment of salinity intrusion in the James and Chickahominy
12 rivers as a result of simulated sea-level rise in Chesapeake Bay, East Coast, USA. *Journal of*
13
14 903 *Environmental Management*, **111**, 61–69. <http://dx.doi.org/10.1016/j.jenvman.2012.06.036>.
15 904
- 16
17 905 Ross R.C., Najjar R.G., Li M., Mann M.E., Ford S.E. & Katz B. 2015 Sea-level rise and other influences
18 on decadal-scale salinity variability in a coastal plain estuary. *Estuarine, Coastal and Shelf*
19 906 *Science*, **157**, 79–92. <http://dx.doi.org/10.1016/j.ecss.2015.01.022>.
20 907
- 21
22 908 Sanders B.F. & Piasecki M. 2002 Mitigation of salinity intrusion in well-mixed estuaries by optimization
23 of freshwater diversion rates. *Journal of Hydraulic Engineering*, **128**(1), 64–77.
24 909
25
26 910 [http://dx.doi.org/10.1061/\(ASCE\)0733-9429\(2002\)128:1\(64\)](http://dx.doi.org/10.1061/(ASCE)0733-9429(2002)128:1(64)).
- 27
28
29 911 Scholz M. 2014 Rapid assessment system based on ecosystem services for retrofitting of sustainable
30 drainage systems. *Environmental Technology*, **35**(9–12), 1286–1295.
31 912
- 32
33 913 Shaha D.C., Cho Y.K., Kwak M.T., Kundu S.R. & Yung K.T. 2011 Spatial variation of the longitudinal
34 dispersion coefficient in an estuary. *Hydrology and Earth System Sciences*, **15**, 3679–3688.
35 914
- 36
37 915 Shanno D. F. & Phua K. H. 1980 Remark on algorithm 500, a variable metric method for unconstrained
38 minimization. *ACM Transaction on Mathematical Software*, **6**(4), 618–622.
39 916
- 40
41 917 Sierra J.P., Sánchez-Arcilla A., Figueras P.A., González Del Río J., Rassmussen E.K. & Mösso C. 2004
42 Effects of discharge reductions on salt wedge dynamics of the Ebro River. *River Research and*
43 918 *Applications*, **20**(1), 61–77. <http://dx.doi.org/10.1002/rra.721>.
44 919
- 45
46 920 Trieu T.T.N. & Phong N.T. 2015 The impact of climate change on salinity intrusion and Pangasius
47 (Pangasianodon Hypophthalmus) farming in the Mekong Delta, Vietnam. *Aquaculture*
48 921 *International*, **23**, 523–534. <http://dx.doi.org/10.1007/s10499-014-9833-z>.
49 922
- 50
51 923 Turrell W.R., Brown J. & Simpson J.H. 1996 Salt intrusion and secondary flow in a shallow, well-mixed
52 estuary. *Estuarine, Coastal and Shelf Science*, **42**(2), 153–169.
53 924
54
55 925 <http://dx.doi.org/10.1006/ecss.1996.0012>.
56 926

- 926 Uncles R.J. & Stephens J.A. 1996 Salt intrusion in the Tweed estuary. *Estuarine, Coastal and Shelf*
1
2 927 *Sciences*, **43**(3), 271–293. <http://dx.doi:10.1006/ecss.1996.0070>.U.S.
- 3
4 928 Climate Change Science Program, 2009. Coastal Sensitivity to Sea-level Rise: a Focus on the Mid-
5
6 929 Atlantic Region. A Report by the U.S. Climate Change Science Program and the Subcommittee on
7
8 930 Global Change Research. [James G. Titus (Coordinating Lead Author), K. Eric Anderson, Donald
9
10 931 R. Cahoon, Dean B. Gesch, Stephen K. Gill, Benjamin T. Gutierrez, E. Robert Theiler & S.
11
12 932 Jeffress Williams (Lead Authors)]. U.S. Environmental Protection Agency, Washington, D.C.,
13
14 933 USA, 320 pp. Available on-line at: [http://www.epa.gov/climatechange/effects/](http://www.epa.gov/climatechange/effects/coastal/sap4-) coastal/sap4-
15
16 934 1.html.
- 17
18 935 Woodworth P.L., Teferle N., Bingley R., Shennan I. & Williams S.D.P. 2009 Trends in UK mean sea
19
20 936 level revisited. *Geophysical Journal International*, **176**(22), 19–30. <http://dx.doi:10.1111/j.1365->
21
22 937 246X.2008.03942.x.
- 23
24 938 Xiao H., Huang W., Johnson E., Lou S. & Wan W. 2014 Effects of sea level rise on salinity intrusion in
25
26 939 St. Marks River Estuary, Florida, USA. *Journal of Coastal Research*, **68**(special issue), 89–96.
27
28 940 <http://dx.doi.org/10.2112/SI68-012.1>.
- 29
30 941 Xu Y., Zhang W., Chen X., Zheng J., Chen X. & Wu H. 2015 Comparison of analytical solutions for salt
31
32 942 intrusion applied to the Modaomen Estuary. *Journal of Coastal Research*, **31**(3), 735–
33
34 943 741. <http://dx.doi:10.2112/JCOASTRES-D-14-00193>.
- 35
36 944 Xue P., Chen C., Ding P., Beardsley R.C., Lin H., Ge J. & Kong Y. 2009 Saltwater intrusion into the
37
38 945 Changjiang River: a model-guided mechanism study. *Journal of Geophysical Research*, **114**(2),
39
40 946 1–15. <http://dx.doi:10.1029/2008JC004831>.
- 41
42 947 Yang Z., Wang T., Voisin N. & Copping A. 2015 Estuarine response to river flow and sea-level rise
43
44 948 under future climate change and human development. *Estuarine and Coastal and Shelf Science*,
45
46 949 **156**, 19–30. <http://dx.doi:10.1016/j.ecss.2014.08.015>.
- 47
48 950 Zhang Z., Cui B., Ou B. & Fan X. 2012 Wetland network design for mitigation of saltwater intrusion by
49
50 951 transferring tidal discharge. *Clean – Soil, Air, Water*, **40**(10), 1057–1063.
- 51
52
53
54
55
56
57
58
59
60
61
62
63
64
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Table 1. References related to salinity intrusion and sea level rise (SLR)

Reference	Location	Finding
Kurup et al. (1998)	Swan River estuary, Australia	The freshwater river inflow is the most critical process impacting on the salt wedge site.
Liu et al. (2004)	Tanshui river system, Taiwan	The average annual salinity was 8.5 parts per thousand (ppt) and 12.8 ppt before and after the construction of the reservoir in this order.
Liu (2005)	Keelung River, Taiwan	The boundaries of saline intrusion before regulation of the channel were further upstream of the river compared to after channel regulation.
Liu et al. (2007)	Danshuei River estuarine, Taiwan	The two-layered circulation within the estuary prevails most often at the estuary mouth.
Xue et al. (2009)	Changjiang River, China	The salt water movement resulted from a complex non-linear interaction process in relation to freshwater flow upstream, tidal currents, water mixing, wind impact and saltwater distribution.
Jeong et al. (2010)	Modaomen Estuary, China	During neap tides, the estuary gains salt, whilst during spring tides, it loses salt.
Cai et al. (2012)	Modaomen Estuary, China	The estuary would be increasingly exposed to salt intrusion and to flooding from storm surges as a result of further deepening, which facilitates the penetration of storm surges into the system.
Bunyan and Dutta (2012)	Goral river network, Bangladesh	A 59-cm rise in sea level increases the salinity by 0.9 ppt at a distance of 80 km upstream of the river mouth.
Rice et al. (2012)	Chesapeake Bay, USA	With 50 cm and 100 cm sea level rises, the maximum salinity increases would be 2 and 4 ppt, respectively.
Chen et al. (2013)	Wu River estuary, Taiwan	More tidal energy will spread into the estuary after weir construction.
Mendes et al. (2013)	Douro estuary, Portugal	The amplitude of the principal lunar semi-diurnal tide constituent increased for almost the entire estuarine area
Liu and Liu (2014)	Wu River estuary, Taiwan	The overall flushing time extracted from the system for high flow without sea level rise is lower in comparison to the sea level rise.
Chen et al. (2014)	Tamsui River estuarine, Taiwan	The maximum increment of depth-averaged and tidal-averaged salinity would be 1.1, 2.4 and 3.0 psu in this order for corresponding sea level rises of 0.34, 1.05 and 1.40 m at the middle region of the estuary under the average discharge scenario.
Kuang et al. (2014)	Yangtze River Estuary	The ebb flow split ratio increases up to almost 5% under a 2-m sea level rise scenario.
Liu and Liu (2014)	Wu River estuary, Taiwan	The tidal expedition further upstream increases by 500 and 900 m under the Q10 and Q90 discharge scenarios, respectively.
Xiao et al. (2014)	Marks River Estuary, USA	A 0.85-m sea level rise caused a substantial increase in salinity near the Wakulla River with an increase of 9.2 ppt for the surface and 12.7 ppt for the bottom salinity.
Chua and Xu (2014)	No specific location (i.e.: idealised estuary model)	A sea level rise causes a strong longitudinal saline water gradient, higher longitudinal dispersion coefficients and increased saline water movement.

Table 2. Most widely used salinity intrusion models

Model	Basic model features	Practical salinity intrusion applications
Statistical	-	Gibson and Najjar (2000); Hilton et al. (2008); Zhang et al. (2010); Cai et al. (2012)
Linear programming	-	Sanders and Piasecki (2002)
1D ^a	-	Bunyan and Dutta (2012); Fleenor and Bombardelli (2013)
2D ^b	Laterally averaged	Kurup et al. (1998); Hsu et al. (1999);
	Vertically averaged	Liu et al.(2001); Liu et al. (2004); Liu (2005)
	MIKE 21 ^d	Kuang et al. (2014)
	MIKE 11 ^e	Dat et al. (2011)
3D ^c	EFDC ^f	Jeong et al. (2010); Gong and Shen (2011); Rice et al. (2012); Hong and Shen (2012); Qiu and Zhu (2013)
	SELFE ^g	Chen et al. (2013); Liu and Liu (2014)
	Un-TRIM ^h	Liu et al. (2007)
	FVCOM ⁱ	Xue et al. (2009)
	MHIOD ^j	Mendes et al. (2013)
	POM ^k	Xiao et al. (2014)
	SUNTANS ^l	
	Chua and Xu (2014)	

^aOne dimensional.

^bTwo dimensional.

^cThree dimensional.

^dTwo-dimensional Hydrodynamic Model.

^eTwo-dimensional Hydrodynamic Model.

^fEnvironment Fluid Dynamics Code.

^gEulerian-Lagrangian Finite-Element.

^hA three-dimensional, time-dependent, baroclinic, hydrodynamic and salinity numerical model.

ⁱFinite Volume Coastal Ocean Model.

^jA baroclinic finite volume two-dimensional numerical model.

^kPrinceton Ocean Model.

^lStanford Unstructured Nonhydrostatic Terrain-following Adaptive Navier–Stokes Simulator.

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Figure 1

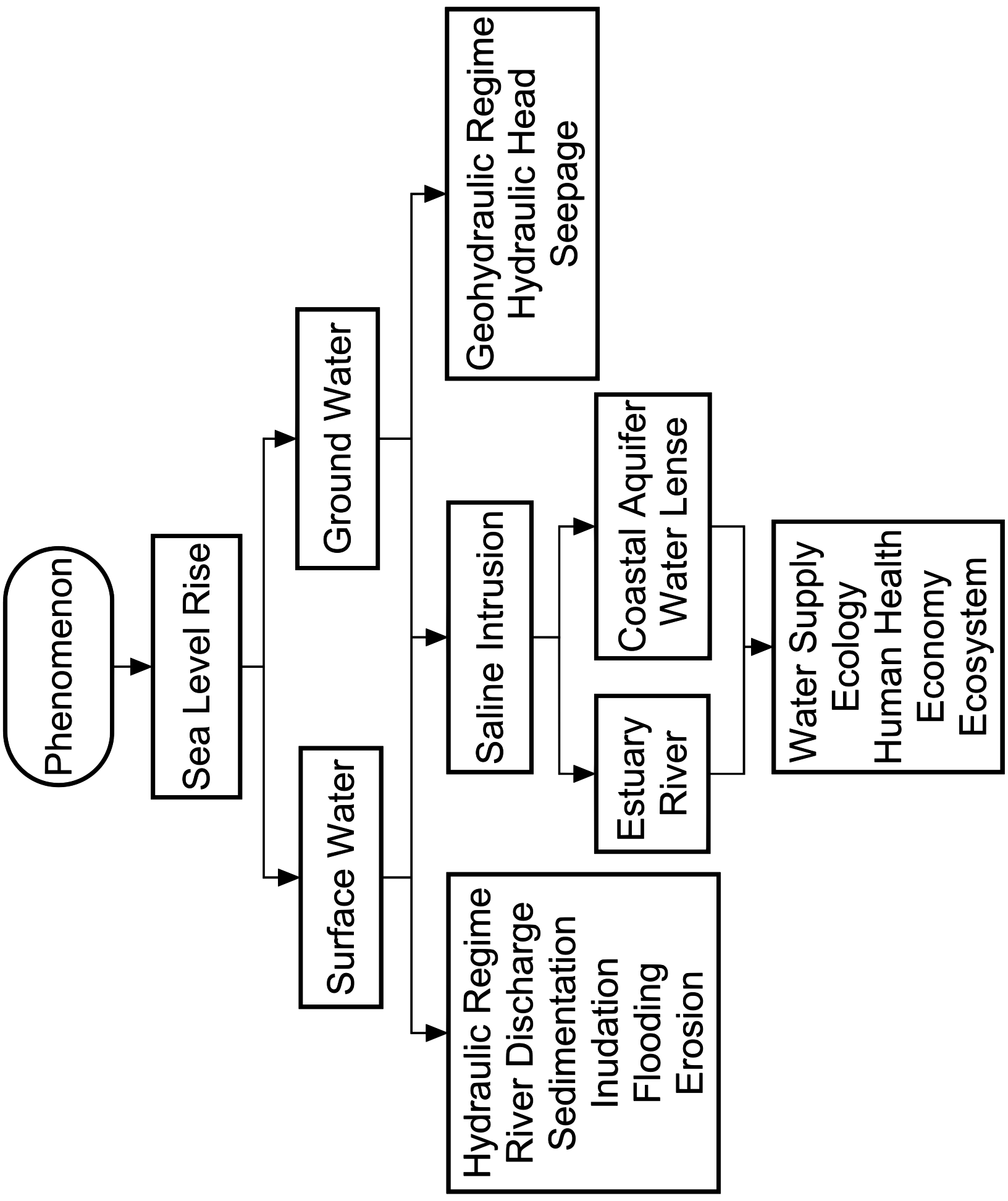


Figure 2

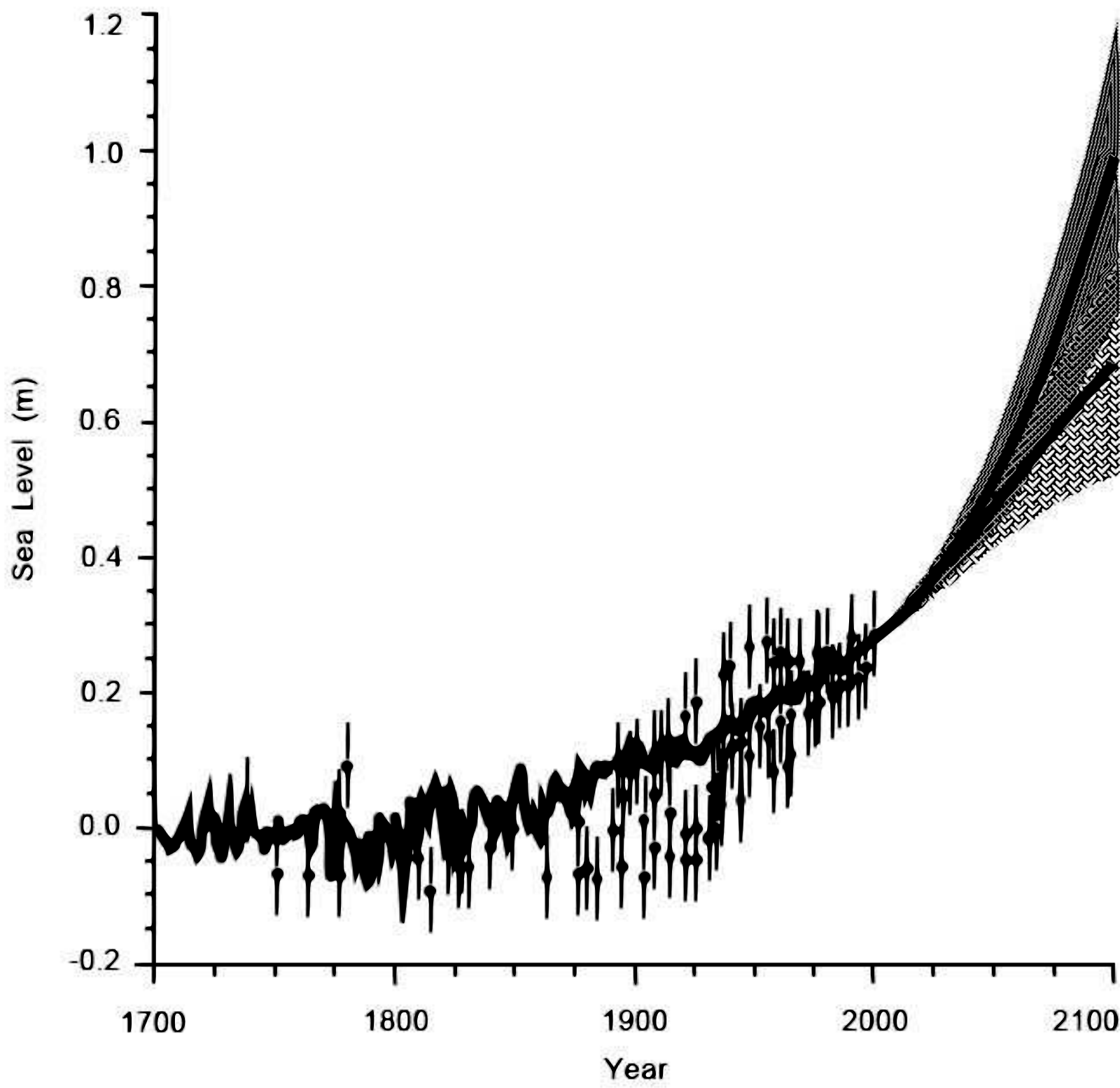


Figure 3

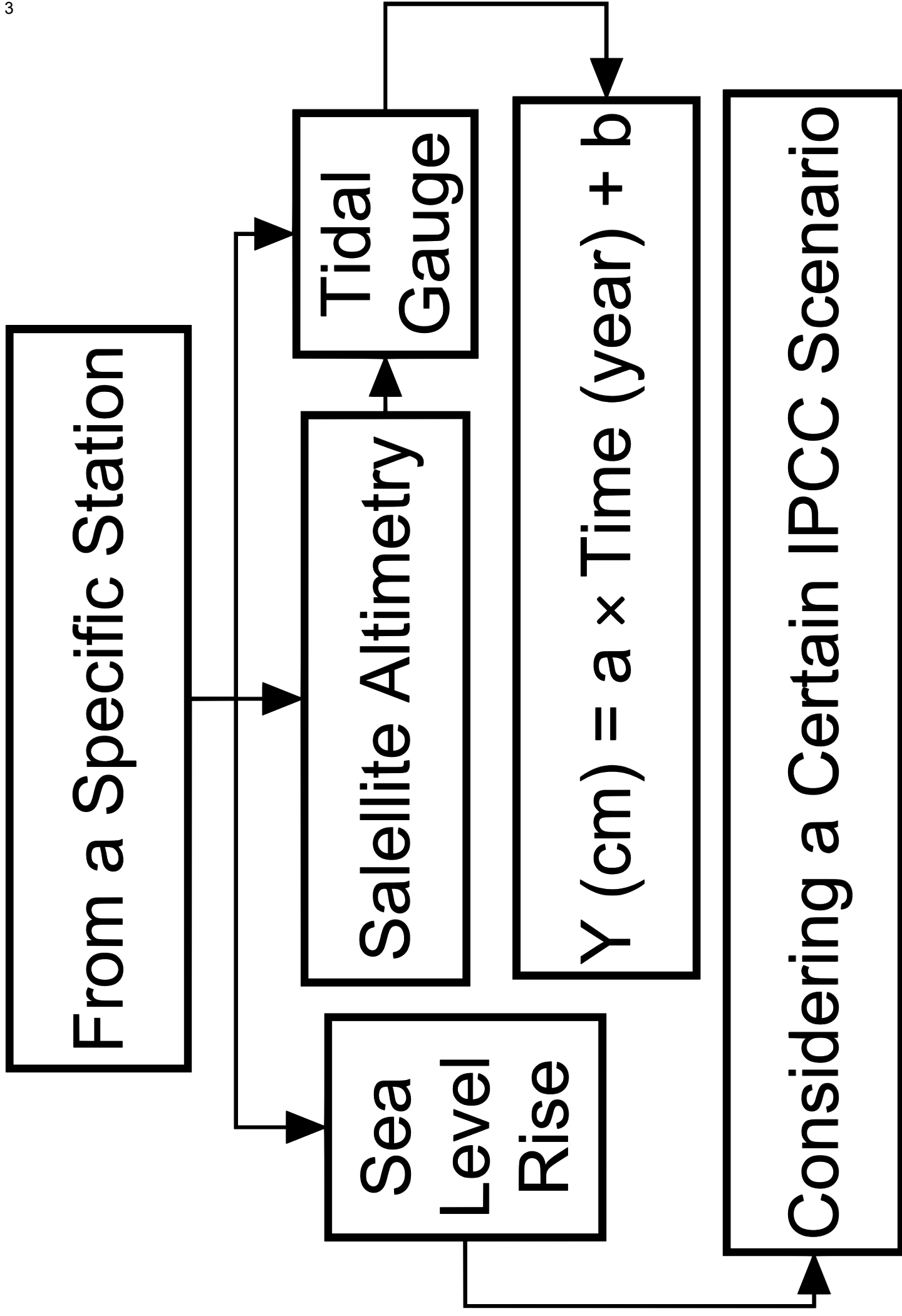


Figure 4

Input Data:
Tidal Elevation and Stream Flow



Hydrodynamic Model



With SLR

Without SLR



Flow/Velocity



Solute Transport Model



With SLR

Without SLR



Saline Water Length and
Saline Water Concentration