

Short-term biotic removal of dissolved organic nitrogen (DON) compounds from soil solution and subsequent mineralisation in contrasting grassland soils

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1	Short-term biotic removal of dissolved organic nitrogen (DON) compounds from soil		
2	solution and subsequent mineralisation in contrasting grassland soils		
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13 ABSTRACT

Cycling of low molecular weight dissolved organic nitrogen compounds constitutes an 14 important component of soil organic matter turnover in soils. Here we determined how rapidly 15 16 grassland soils can cycle urea, compared to the amino acid L-alanine, and the peptide Ltrialanine. Using naturally occurring concentrations of ¹⁴C-labelled compounds the rates of 17 removal from soil solution and subsequent mineralisation were measured. Biotic removal of 18 all three compounds and subsequent mineralisation to CO2 occurred within minutes. This 19 research has demonstrated, for the first time, the potential for rapid removal of urea at low 20 21 concentrations by the soil microbial biomass.

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23 Keywords: Dissolved organic matter, Nutrient cycling, Urine patch, Urea

Adverse ecosystem effects of excess nitrogen (N) have been observed globally (Vitousek et al., 2009). Excess N in grasslands, prone to leaching and N emissions, is typically derived from N amendments including organic manures, urine patches and excessive use of synthetic fertilisers. Of these, urea has frequently been examined due to its importance as a fertiliser (IFA, 2014), and its presence in manures and urine (Ball and Ryden, 1984).

Urea is a low molecular weight dissolved organic N (LMW-DON) compound with a 29 30 C:N molar ratio of 1:2.33, and similar to nitrate and ammonium, is capable of being taken up directly by both plants and microorganisms (Berman and Bronk, 2003; Wang et al., 2008). The 31 32 extent to which plants can acquire LMW-DON and the degree to which it leaches down the soil profile, is critically dependent on the activity of the soil microbial biomass (SMB; Jones et 33 al., 2013). Recent studies indicate that uptake of LMW-DON by the SMB is frequently driven 34 35 by carbon (C) demand rather than N (Farrell et al., 2014). Therefore, the presence of C within 36 urea may drive its rate of removal from grassland soil solutions. Although urease activity in soil (Nielsen et al., 1998; Bolado-Rodríguez et al., 2005), and to a lesser extent urea 37 38 assimilation by the SMB (Smith et al., 2007) have been investigated, urea removal from the soil solution by the SMB over short time-scales has not. 39

Here SMB removal of ¹⁴C-urea from the soil solution and its subsequent catabolic and
anabolic partitioning was examined in three grassland soils. Microbial cycling of urea was
directly compared to that of other typical LMW-DON compounds found in soils: the amino
acid ¹⁴C-L-alanine and the oligopeptide ¹⁴C-L-trialanine, whose turnover have been extensively
characterised and have also been implicated in direct plant LMW-DON acquisition (Hill et al.,
2011; Wilkinson et al., 2014).

Soil was collected from three separate grazed grassland sites in the UK (Table 1). All
soils were collected towards the end of the growing season (October), with three independent

replicates collected for each type. Soil cores (10×8.5 cm; $h \times i.d.$) were kept intact, at fieldmoisture, in gas-permeable bags, in the dark at 4°C prior to use.

To characterise LWM-DON in each soil, porewater was obtained from intact soil cores, 50 with the root mat removed, by centrifugation-drainage (Giesler and Lundstöm, 1993). Soluble 51 N was determined as described by Farrell et al. (2013) and Sullivan and Havlin (1991). All 52 experimentation with ¹⁴C-labelled compounds was performed on < 2 mm sieved soil from 53 separately taken soil samples, which had equilibrated to 20°C overnight. The rate of LMW-54 DON depletion from soil solution was measured according to Hill et al. (2008). Briefly, 1 g 55 56 soil (dry weight equivalent; DW) was placed in a microcentrifuge tube with a hole pierced in the bottom. This was placed inside another microcentrifuge tube. 300 µl of either ¹⁴C-labelled 57 urea, L-alanine or L-trialanine (10 μ M, 0.9 kBq mL⁻¹) was then applied to the soil surface and 58 59 allowed to infiltrate the soil (< 2.5 sec, 20°C; associated soil water content increase to 45-52%). This concentration was chosen to reflect the urea and free amino acid concentrations naturally 60 occurring within soil solution (Table 1). At 1, 5, 10, 30 and 60 min after substrate addition, the 61 62 soil was centrifuged (4000 g, 1 min, 4°C; data presented for 0 min in Figures 1 and 2 are assumed) allowing the soil solution to pass to the lower microcentrifuge tube. The ¹⁴C content 63 of the recovered soil solution was determined after addition of Scintisafe3 scintillation cocktail 64 (Fisher Scientific, Loughborough, UK) using a Wallac 1404 liquid scintillation counter (Perkin 65 Elmer Life Sciences, Boston, MA). To assess the mineralisation rate of LMW-DON 66 67 compounds, 1 g soil was placed in a glass tube through which air was passed before being transferred through 2 successive 0.1 M NaOH traps to capture evolved ¹⁴CO₂. At 1, 5, 10, 30 68 and 60 min after ¹⁴C-labelled substrate addition (as above), NaOH was replaced and its ¹⁴C 69 content determined as above. To separate biotic (e.g. microbial, enzymatic) and abiotic (e.g. 70 sorption) LMW-DON removal processes, the soil solution recovery experiment was also 71 performed on sterilised soil (autoclaved at 121°C, 20 min). Recovery of ¹⁴C-labelled 72

compounds from the sterilised soil solutions was used to calculate the theoretical maximum
 ¹⁴C-activity (Hill et al., 2008) that could be recovered following complete mixing of amended
 ¹⁴C-labelled treatments with native soil solution. A two-way ANOVA was used to test for
 differences and interactions between soils and LMW-DON treatments.

Complete mixing with native soil solution was not achieved, and after 60 min deviated 77 between 101-153%. Greater than 100% recovery was achieved at all incubation periods, thus 78 demonstrating that no retention of ¹⁴C-compounds occurred in the sterile soils (Wilkinson et 79 al., 2014; see supplementary information for equations), consequently no evidence of abiotic 80 81 loss pathways was observed. However autoclaving soils can increase the solubilisation of soil organic matter (SOM; Powlson and Jenkinson, 1976), which may block adsorption sites that 82 would be available in the living soils. Although soil sterilisation via autoclaving has been found 83 84 to be more effective that CHCl₃ fumigation or gamma irradiation at reducing viable cell 85 numbers (Blankinship et al., 2014).

All ¹⁴C-labelled LMW-DON compounds were rapidly removed from the soil solution 86 (Fig. 1). After 60 min, removal of ¹⁴C-L-alanine and ¹⁴C-L-trialanine was almost complete, at 87 98.7 and 99.5% respectively. However, removal of ¹⁴C-urea was consistently lower at all 88 incubation periods, and after 60 min was 88.7%. Removal from soil solution followed the 89 series: alanine > trialanine > urea (p < 0.001). Across all soils, the half-life of urea, alanine and 90 trialanine in solution was 4.15 ± 0.69 , 0.30 ± 0.04 , 0.94 ± 0.13 min (mean \pm SEM; based on 91 92 fitting first order single exponential decay to the data), respectively. In contrast, no effect of soils was observed on substrate depletion. This is perhaps unsurprising as cycling of key LMW-93 DON compounds can be remarkably similar across diverse soils and systems (Jones et al., 94 95 2009). Slower uptake of urea relative to the other compounds could be linked to lower transporter expression and affinity within the SMB. It has also been proposed that soil 96 microorganisms exist in a C-starved state (Hobbie and Hobbie, 2013). Accordingly, in the soils 97

98 examined here the SMB may have exhibited a preference for compounds with a greater C
99 content and which are commonly present in soil solution (amino acids and peptides are present
100 in rhizodeposits and via protease action on SOM).

101 In contrast to removal from soil solution, soil affected (p < 0.001) mineralisation of different compounds at 1, 5, 10 and 30 min, but not after 60 min (Fig. 2). The differences 102 between soils may, in part, be due to differences in initial soil water content, which can impact 103 net mineralisation rates (Paul et al., 2003). However, the experiment was performed on field 104 moist soils to represent the same preceding climatic conditions rather than target a specific soil 105 water content (Table 1). ¹⁴C-urea had the highest mineralisation rates of the three LMW-DON 106 compounds. This may be attributed to the alternative mineralisation pathway urea can take via 107 108 the enzyme urease, which is encountered both intra- and extracellularly. Extracellular urease 109 has been shown to account for an average of 46% of total urease activity in a range of soils (Klose and Tabatabai, 1999). This may account for the more rapid mineralisation of urea 110 relative to alanine and trialanine. Although ¹⁴C-urea was most rapidly mineralised of the three 111 compounds, only 40-45% of ¹⁴C-urea removed from the soil solution was subsequently respired 112 as ¹⁴CO₂ over 60 min, suggesting that the remaining ¹⁴C-urea was assimilated by the SMB 113 (Nielsen et al., 2008). Another intracellular urea pathway is via ATP: urea amidolyase (Cheng 114 et al., 2005; Strope et al., 2011), which produces NH₃ and HCO₃⁻ via two enzyme (urea 115 carboxylase and allophanate hydrolase) reactions, making it a likely path for urea assimilation 116 117 in soils.

Although it is widely acknowledged that urea is rapidly mineralised in soils, this is the first time that such rapid removal of urea from the soil solution by the SMB has been reported. This suggests that there will be strong microbial competition for urea in soil which may limit its capture by plant roots when present in low concentrations. The fate of urea-derived NH₄⁺

- requires further investigation. Assimilation of urea by the SMB at higher concentrations of ureaand following application of solid urea as fertiliser requires further work.
- 124

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- 128

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	Dystric Cambisol*	Stagni-vertic Cambisol*	Eutric Cambisol**
Sample location	50°46'N, 3°55'W	50°47'N, 3°57'W	53°14'N, 4°01'W
Soil texture	Loam*	Heavy clay*	Sandy clay loam**
Soil pH	5.2 ± 0.07	5.3 ± 0.07	4.8 ± 0.03
Total soil C (g kg ⁻¹ DW)	21.1 ± 0.04	29.1 ± 0.29	27.1 ± 0.19
Total soil N (g kg ⁻¹ DW)	3.6 ± 0.02	4.0 ± 0.02	4.1 ± 0.04
Soil water (kg kg ⁻¹ DW)	0.19 ± 0.01	0.35 ± 0.02	0.29 ± 0.00
Soil solution free amino acids (μM)	11.3 ± 1.08	6.50 ± 1.18	8.00 ± 1.48
Soil solution short peptides (<1 kDa; μ M)	153 ± 47.3	145 ± 24.7	164 ± 52.0
Soil solution NO ₃ -N (mg N l ⁻¹)	10.6 ± 0.92	0.74 ± 0.23	2.22 ± 1.21
Soil solution urea (µM)	$6.22 \hspace{0.1 cm} \pm \hspace{0.1 cm} 0.20 \hspace{0.1 cm}$	7.24 ± 0.62	15.33 ± 8.12
Soil solution NH4-N (mg N l ⁻¹)	0.14 ± 0.02	0.30 ± 0.07	0.23 ± 0.05
Soil solution DOC (mg C l ⁻¹)	29.38 ± 3.25	31.80 ± 5.51	19.98 ± 2.77
Soil solution DON (mg N L ⁻¹)	1.94 ± 0.38	2.17 ± 0.38	1.54 ± 1.26
Soil respiration (mg C kg ⁻¹ dry soil h ⁻¹)	0.37 ± 0.05	0.86 ± 0.32	0.60 ± 0.14

Table 1 Characteristics of the three grassland soils (upper 10 cm) used in the study. Values represent mean \pm SEM (n = 3).

Data gained from the literature are marked with either a * (described by Harrod and Hogan, 2008) or a ** (described by Hill et al., 2012).

201 Figure legends

- 202 Fig. 1. Microbially-mediated depletion of ¹⁴C-labelled alanine, trialanine or urea from soil
- solution in three grassland soils. Data points represent means \pm SEM (n = 3).

204

- **Fig. 2.** Time-dependent cumulative mineralisation of ¹⁴C-labelled alanine, trialanine or urea
- in three grassland soils. Data points represent means \pm SEM (n = 3).