

1 **Multi-Model Skill Assessment of Seasonal**
2 **Temperature and Precipitation Forecasts over**
3 **Europe**

4 **Niti Mishra · Chloé Prodhomme ·**
5 **Virginie Guemas**

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8 **Abstract** There is now a wide range of forecasts and observations of seasonal
9 climatic conditions that can be used across a range of application sectors, in-
10 cluding hydrological risk forecasting, planning and management. As we rely
11 more on seasonal climate forecasts, it becomes essential to also assess its qual-
12 ity to ensure its intended use. In this study, we provide the most comprehen-
13 sive assessment of seasonal temperature and precipitation ensemble forecasts
14 of the EUROSIP multi-model forecasting system over Europe. The forecasts
15 from the four individual climate models within the EUROSIP are assessed
16 using both deterministic and probabilistic approaches. One equally and two
17 unequally Weighted Multi-Models (WMMs) are also constructed from the in-
18 dividual models, for both climate variables, and their respective forecasts are
19 also assessed.

20 Consistent with existing literature, we find limited seasonal climate predic-
21 tion skill over Europe. A simple equally WMM system performs better than
22 both unequally WMM combination systems. However, the equally WMM sys-
23 tem does not always outperform the single best model within the EUROSIP
24 multi-model. Based on the results, it is recommended to assess seasonal tem-
25 perature and precipitation forecast of individual climate models as well as their
26 multi-model mean for a comprehensive overview of the forecast skill.

27 **Keywords** Seasonal Climate Forecast · Probability Ensemble Forecast ·
28 Weighted Multi-Model · Forecast Verification

Department of Earth Sciences, Barcelona Supercomputing Centre
Carrer de Jordi Girona, 29-31, 08034 Barcelona, Spain
E-mail: niti.mishra@bsc.es
E-mail: chloe.prodhomme@bsc.es
E-mail: virginie.guemas@bsc.es

1 Introduction

Recent hydrological extreme events demonstrate the vulnerability of European society to water-related natural hazards and there is a strong evidence that climate change will worsen these events (Lavell et al. 2012, chap1; National Academies of Sciences and Medicine 2016). The impacts of these hydrological extreme events can be reduced by early-warning design support systems (Wanders and Wood 2016). Hydrological simulations in these support systems rely on initial land surface conditions from upstream river flow, snow cover, soil moisture and/or skillful seasonal prediction of continental meteorological conditions, such as temperature and precipitation (Wanders and Wada 2015; Yuan et al. 2016). The predictability of precipitation and temperature is exploited particularly for long-term hydrological forecasts (Velázquez et al. 2009; Yuan et al. 2016) and thus, high-quality Seasonal Climate Forecasts (SCFs) are essential for the success of seasonal hydrological forecasting based on climate models.

SCFs are forecasts of climate conditions at timescales of a few weeks up to a few months, for statistics such as monthly/seasonal averages of temperature and/or precipitation or frequency of occurrences of extreme events. SCFs are possible due to the long-term predictability of the oceanic circulation (i.e. up to a few years) and by the fact that the variability in tropical Sea Surface Temperature (SST) has a significant global impact on the atmospheric circulation (Balmaseda and Anderson 2009; Doblus-Reyes et al. 2013). Considerable efforts have been made in the field to better represent the coupled ocean-atmospheric dynamics and to improve the operational Climate Forecast Systems (CFSs) such as the National Centers for Environmental Prediction (NCEP; Saha et al., 2014) and the Predictive Ocean Atmosphere Model for Australia (POAMA; Colman, 2005), which are single-model CFSs as well as the European Multimodel Seasonal to Interannual Prediction (EUROSIP; Vitart et al. 2007; EUROSIP 2016; Stockdale 2013) system, which comprises of four independent CFSs.

Generally, forecast skill of seasonal climatic conditions in areas influenced by ENSO is higher than in the extra-tropical regions (Alexander et al. 2002; Kumar et al. 2013; Palmer et al. 2004; Sordo et al. 2008). In Europe, stratospheric processes (Bell et al. 2009), snow cover (Senan et al. 2016), soil moisture (Prodhomme et al. 2016) and sea-ice (Guemas et al. 2016) are also proven to be effective sources of predictability. Recently, the North Atlantic Oscillation (NAO) has also been reported as an important source of predictability for European winter climate (Athanasiadis et al. 2017; Scaife et al. 2014). Yet, the overall seasonal forecast skill over Europe for surface variables is still quite low (Arribas et al. 2011; Kim et al. 2012; Scaife et al. 2014).

An ensemble forecast is a set of forecasts that generate a range of future climate possibilities. Ensemble forecasts are often preferred over deterministic ones because they can convey the uncertainties that arise due to the inability to accurately model atmospheric dynamics and the initial condition uncertainty (Hawkins and Sutton 2009, 2011; Lorenz 1963; Palmer et al. 2004; Tebaldi

74 and Knutti 2007). To obtain ensemble forecast, a climate model is run multi-
75 ple times, each time with slightly different initial conditions and with slightly
76 perturbed numerical models. Each forecast in an ensemble, known as a mem-
77 ber, are then used to calculate the probability distribution of the potential
78 near-term future climate (Bröcker and Smith 2008; Fortin et al. 2006; Wilks
79 2006).

80 In addition to ensemble prediction, combining ensembles of multiple CFSs
81 to make climate predictions has also garnered a lot of attention (Doblas-Reyes
82 et al. 2003; Palmer et al. 2005; Rodrigues et al. 2014; Weigel et al. 2010; Yun
83 et al. 2003). The EUROSIP multi-model, which became operational in 2005, is
84 the result of DEMETER (Palmer et al. 2004) and other research projects that
85 confirmed the scientific benefits of combining forecasts from several climate
86 models. Such multi-model predictions address issues of structural uncertainties
87 within models that arise due to incomplete physical parameterizations and nu-
88 merical approximations (Palmer et al. 2004). In general, equally weighting each
89 of the CFSs has been recognized to have consistently better performance than
90 that of the individual models (DelSole 2007; Hagedorn et al. 2005; Kharin and
91 Zwiers 2002; Peng et al. 2002). This is because random individual model errors
92 tend to compensate one another and the robust predictable signal tend to stand
93 out by averaging across a number of models. This is particularly important
94 for medium-to-long range forecasting, where the timescale over which model
95 errors accumulate are much longer and can significantly degrade long-term
96 forecasts. On the other hand, improved performance of unequally Weighted
97 Multi-Model (WMM) systems have also been reported (Krishnamurti et al.
98 2000; Robertson et al. 2004; Rodrigues et al. 2014; Wanders and Wood 2016).
99 Intuitively, it makes sense to give more (less) weight to forecasts from a model
100 that has consistently better (poor) historical performance. However, consen-
101 sus on the optimal way of weighing the different models has yet to be reached
102 (DelSole et al. 2013; Tebaldi and Knutti 2007).

103 Whether we use deterministic, probabilistic or some weighted combination
104 of forecasts from multiple models, they are ultimately beneficial only if they
105 have skill and can add value to the users. The objective of this study is to assess
106 the seasonal forecasting skill of each of the forecast system in the EUROSIP
107 multi-model and to compare it with that of equally and unequally WMMs in
108 order to provide users in hydrology an overview on current potential and/or
109 limits of seasonal temperature and precipitation predictability over Europe. A
110 systematic investigation across different models of EUROSIP and for different
111 seasons specifically over the European region is still lacking and this study
112 aims to highlight the need for further studies by contributing to the limited
113 extant literature. The assessment is done for winter and summer, temperature
114 and precipitation forecasts over a period of 21 years (1992-2012) in terms of
115 the Anomaly Correlation Coefficient (ACC) for deterministic forecasts and the
116 Continuous Ranked Probability Skill Score (CRPSS) for probabilistic forecasts
117 on each grid point.

118 The article is structured as follows: Section 2 describes the datasets used,
119 the methods applied to assess the forecast skill and to construct the WMMs.

120 Section 3 presents the subsequent results followed by a discussion and a final
121 conclusion in Section 4 and 5, respectively.

122 2 Data and Methods

123 2.1 Data

124 This study relies on a comprehensive set of seasonal temperature and precip-
125 itation forecasts from the EUROSIP multi-model over the European region
126 specified as 20°W-70°E and 25°N-75°N, for the period 1992-2012. Four indi-
127 vidual CFSs – Global Seasonal forecasting system version 5 (GloSea5) from
128 Met Office, System4 of European Center for Medium Range Weather Fore-
129 casts (ECMWF), System2 of NCEP and System5 of Meteo France (MF), are
130 integrated into one common EUROSIP multi-model framework (EUROSIP
131 2016). These are the common choice of operational multi-model in the Eu-
132 ropean region (Soares and Dessai 2015; Stockdale 2013) and we select the
133 longest available hindcast period in common for this study. The number of
134 ensemble members and the horizontal resolutions of these four climate models
135 are given in Table 1. More details on the dynamical cores and the physical pa-
136 rameterizations of individual models within EUROSIP can be found in their
137 corresponding documentations (MacLachlan et al. 2015; Molteni et al. 2011;
138 Saha et al. 2014; Volodire et al. 2013).

139 [Table 1 about here.]

140 The reference dataset for temperature is obtained from the ERA-Interim
141 (ERAINT) database, which includes a 4D variational analysis with a 12-hour
142 analysis window (Dee et al. 2011). The spatial resolution of the dataset is \cong
143 80 km (T255 spectral) on a reduced Gaussian grid with 60 vertical levels from
144 the surface up to 0.1 hPa (Dee et al. 2011). The results (not presented) are
145 insensitive to the comparison with the observation dataset from Global His-
146 torical Climatological Network (GHCN 2.2). For precipitation, the reference
147 dataset is provided by the Global Precipitation Climatology Project (GPCP),
148 which comprises a gridded analysis based on gauge measurements and satellite
149 estimates of precipitation (Adler et al. 2003).

150 The original values of both forecasts and observations are interpolated
151 using a bilinear interpolation to match the coarsest grid among each climate
152 variable. The coarsest grid is chosen as a preferred grid for such interpolation
153 method (Starks et al. 2003). All computations are done on grid point by grid
154 point basis. The sea points are masked and only data over land is assessed.

155 The forecasts from the four models and the reference datasets are available
156 with monthly averages of daily mean temperature and precipitation values
157 for the period 1992-2012. The study is performed at seasonal (average of three
158 months) timescale for winter and summer seasons. Winter consists of forecasts
159 from December to February (DJF) while summer consists of forecasts from
160 June to August (JJA). All forecasts are initialized around the first day of the

161 month preceding the target season. These particular seasons and years are
162 selected for study because a homogeneous history of hindcasts on a monthly
163 timescale across all four participating forecast systems is available only for
164 this time period. This is an important limitation of the study because a short
165 time series of 21 years usually cannot accurately account for the sensitivity
166 of climate system performance to the chaotic nature of climate, which differs
167 greatly within the various regions of Europe. The data limitation also extends
168 to the verification metrics used and the methodologies applied to combine
169 forecasts as statistics derived from limited number of data is expected to suffer
170 from uncertainty due to sampling error. Longer common period hindcasts are
171 essential in studies to allow the results of analysis to be extended.

172 All calculations in this study are applied to the forecast anomalies com-
173 puted with respect to model's own climatology. Therefore, the ability to predict
174 departures from the seasonal cycle is measured rather than the absolute values
175 of temperature and precipitation. Thus, the model bias does not appear in the
176 verification metrics (or only indirectly since it might affect the variability).

177 2.2 Methods of Verification

178 2.2.1 Anomaly Correlation Coefficient (ACC)

179 In this study, for all deterministic forecasts i.e. the ensemble mean, ACC is
180 used to assess the forecast skill. ACC is the most widely used skill metric for
181 SCF quality (Doblas-Reyes et al. 2013; Fricker et al. 2013; Scaife et al. 2014),
182 due to its invariant property (i.e. not affected by certain data transformation).
183 ACC assesses the degree of linear correspondence between the target forecast
184 anomalies and the anomalies of the observed climate variable. Additionally,
185 for linearly re-calibrated forecasts, the squared ACC is equivalent to the mean
186 squared skill score (Siegert et al. 2017). It is worth noting however, that cor-
187 relation coefficient are extremely noisy in smaller sample sizes, meaning small
188 changes in forecast values can impact correlation skill significantly. We test the
189 significance of ACC at 5% significance level, controlling the False Discovery
190 Rate (FDR) (Benjamini and Hochberg 1995).

191 2.2.2 Continuous Ranked Probability Skill Score (CRPSS)

192 The second metric selected to assess forecast skill is the Continuous Ranked
193 Probability Score (CRPS), which is a standard measure for assessing the ac-
194 curacy and reliability aspects of probabilistic forecasts. CRPS evaluates the
195 predictive skill of the full probability distribution of forecast obtained from
196 the ensemble members (Hersbach 2000; Matheson and Winkler 1976). Such
197 evaluation is desirable since climate forecasts are used as forcings in models
198 such as the hydrological models (Boucher et al. 2009; Candille and Talagrand
199 2005; Gneiting et al. 2005; Murphy 1969).

Given that F is the cumulative density function of ensemble forecasts and y is the value that actually occurred, the CRPS is defined as:

$$CRPS(F, y) = \int_{-\infty}^{\infty} [F(t) - H(t - y)]^2 dt, \quad (1)$$

where $H(t - y)$ denotes the Heaviside function that takes the value of 0 when $t < y$ and 1 otherwise (Hersbach 2000; Matheson and Winkler 1976). Thus, the CRPS measures the difference between the predicted and observed cumulative distributions. For deterministic forecasts, the average CRPS becomes the mean absolute error and therefore, has similar interpretation.

The skill score based on CRPS is CRPSS, computed as:

$$CRPSS = \frac{CRPS_f - CRPS_{clim}}{CRPS_{perf} - CRPS_{clim}}, \quad (2)$$

where $CRPS_f$, $CRPS_{clim}$ and $CRPS_{perf}$ stand for CRPS of the forecast in question, of the reference/benchmark forecast and that of the perfect forecast, respectively. In this study, climatology is used as the reference forecast, which refers to the average conditions over some recent reference period. Skill scores below 0 are unskillful compared to a naïve climatological forecast. Those equal to 0 are no better than that of climatology and anything above 0 (up to 1) signals an improvement upon climatology. The standard deviation of the skill score is approximated by propagation of uncertainty and the significance is measured at 95% confidence interval.

2.2.3 Fair Continuous Ranked Probability Skill Score (FCRPSS)

One drawback of the CRPS is that it inflates the score for models with higher number of ensemble members. To correct for this, Ferro et al., (2014; 2008) recommended the Fair Continuous Ranked Probability Score (FCRPS), which evaluates the underlying ensemble distribution and is independent of the empirical distribution of the ensemble members (Fricker et al. 2013). Results of FCRPSS (skill score based on FCRPS) are also provided.

2.3 Methods for Weighted Multi-Model (WMM) Combination

An important objective of this study is to combine forecasts from dynamical systems to estimate a single optimal forecast with an aim to understand benefits of such combination on the overall forecast quality. Separate models are established for each season and grid cell independently. Three methods of combinations are used in this study:

2.3.1 Multi-Model Mean (MMM)

The first combination approach consists of an equally WMM system, which is obtained by averaging ensembles of each CFS and then again, averaging these four ensemble means to obtain a multi-model ensemble mean anomalies. Hereinafter, this method is referred to as the Multi-Model Mean (MMM). This is one of the most commonly used method to combine forecasts of independent CFSs (DelSole et al. 2013; Kharin and Zwiers 2002; Krishnamurti et al. 2000). The basic idea behind this approach is the assumption that each individual CFS is equally likely to represent the truth whatever its performance (Wanders and Wood 2016).

2.3.2 Best OLS Combination Method (BOCM)

Various forms of regression have been tested on seasonal and weather forecasts to obtain optimal weights based on historical performance of the model (DelSole et al. 2013; Kharin and Zwiers 2002; Rodrigues et al. 2014; Weigel et al. 2008). The second method uses the Ordinary Least Squares (OLS) regression technique to obtain optimal weights. 15 possible OLS models are built out of the ensemble mean of each of the four available CFSs - one Multiple Linear Regression (MLR) model with all four CFSs, four MLR models with only three CFSs, six MLR models with only two CFSs and four linear regression models with only one CFS. For each of these 15 OLS combinations, ensemble mean of the participating CFS(s) are regressed onto the corresponding observations and their respective weights are the regression coefficients estimated from the data. Out of the 15 possible OLS models, the one that has highest correlation with the observation dataset is chosen as the Best OLS Combination Model (BOCM) for each grid point.

2.3.3 Correlation As Weight Method (CAWM)

The final weighted combination method uses as weights the ACC value between ensemble mean anomalies of each CFS and the anomalies of the observation. While correlation does not take into account the system performance in terms of variance, it is often the value relied upon for forecast verification. In addition, correlations are indicative of model performance and thus, it is reasonable to think of correlation values as potentially trustworthy weights. Note that this method may choose a CFS with only a minor correlation improvement among the competing CFSs. Here, the ACC value of each CFS is first multiplied to its respective forecast value. They are then added together and divided by the sum of their ACCs to standardise the forecast value. Hereinafter, this model is referred to as Correlation As Weights Model (CAWM).

2.3.4 Evaluation of Optimal Weights

Historical data is required to not only build statistical models but to also evaluate them. In this study there are only 21 years of records, which is considered not long enough to develop and validate regression-based model. However, a homogeneous history of hindcasts on a monthly timescale across all four participating forecast systems is available only for this time period. This is a major limitation of this study, albeit common in seasonal climate forecasting (Kumar 2009; Shi et al. 2015). To address the issue of small sample size, we apply leave-one-out cross-validation procedure in both WMMs (Efron 1983; Molinaro et al. 2005). This means for each forecast year, the model weights are estimated from the other 20 years of data and a seasonal forecast is made for that year. The process is repeated over each of the 21 years and the resulting hindcasts are then compared with the corresponding observation. An accuracy estimate obtained using leave-one-out cross-validation is known to be almost unbiased but has high variance (Chapelle et al. 2002; Efron 1983).

Another possible reason for unstable weights in linear regressions is the collinearity among the predictors, which can be dealt with by ridge regression (DelSole et al. 2013). However, multicollinearity among EUROSIPs CFSs was found not to be high enough to pursue further (See Fig. S1-2 in supplementary section).

Finally, the weights in both WMMs are constrained meaning both zero or negative weights are not allowed in the model. Model with unconstrained weight was tested by Wanders and Wood (2016) but omitted eventually due to poorer performance. This is because unconstrained models give rise to overconfident estimates of weights when the number of sampling years is small. Besides, it is reasonable to assume constrained weights because a CFS that consistently lacks skill for any given region can be removed from the combined model.

3 Forecast Skill Assessment of EUROSIP

3.1 Assessment of Individual Model Ensemble Mean Anomalies

3.1.1 Strength and Weakness of the Individual Models

This section focuses on the evaluation of the prediction skill of the individual CFSs of EUROSIP in terms of ACC. Figure 1 shows the ACC of seasonal temperature anomalies for both winter and summer seasons. There is a difference in skill exhibited by the four models between the two seasons. GloSea5 has some skill over Europe during winter, although it is not statistically significant. ECMWF has high statistically significant skill over the British Isles, Southern Sweden and parts of Central Europe during winter. NCEP exhibits some statistically significant skill over the North-Eastern Europe (close to Barents

304 sea) and the skill of MF is significantly higher during winter over the Western
305 Europe, the British Isles and the south of Scandinavia.

306 During summer, Glosea5 has notably higher, statistically significant skill
307 over the Northern Scandinavia (close to the Norwegian sea) and the Southern
308 Europe. ECMWF exhibits higher statistically significant skill notably over
309 large parts of the Southern and the Eastern Europe. For NCEP, the summer
310 seasonal temperature skill is higher over the East-Central Europe. MF exhibits
311 limited skill (although not significant) mostly over the Southern regions during
312 summer.

313 [Fig. 1 about here.]

314 [Fig. 2 about here.]

315 Figure 2 shows the results of the same evaluation as that of Figure 1 but for
316 seasonal precipitation. It can be noted that the skill across Europe for seasonal
317 precipitation is very low and sporadic. The skill is higher in winter for Glosea5
318 and ECMWF, mostly concentrated over the Eastern Europe. During winter,
319 NCEP shows some significant skill over the Northern Scandinavia and MF
320 exhibits significant skill over most of Scandinavia and over the British Isles.
321 Higher skill in winter could be because precipitation is hard to both observe
322 and to forecast, given its high variability during the dry months of summer.
323 Conversely, winter precipitation are more dependent on large scale circulation,
324 such as the NAO that has recently shown to have predictability and could be
325 one of the sources of skill here (Scaife et al. 2014; Trigo et al. 2002).

326 For summer, Glosea5 and ECMWF exhibit significant skill over the Mediter-
327 ranean region. NCEP has some significant skill over North-Eastern Europe.
328 The summer seasonal precipitation skill of MF is notably low over Europe.
329 The skill pattern for seasonal precipitation must be considered with caution
330 however, because in regions with limited rain, small changes in observed precip-
331 itation can greatly impact correlation values. Thus, more evidence is needed to
332 make conclusions about EUROSIP's seasonal precipitation forecast skill over
333 Europe.

334 3.1.2 Utilizing Differences of Individual Model Skill

335 An important benefit of using multiple models is their potential capability
336 to complement each other. It is unknown a priori which model performs best
337 in which region. Thus, the different levels of skill of the different models can
338 be exploited in an operational context. Figure 3 shows which model has the
339 highest correlation at each grid point. While MF has an overall relatively low
340 seasonal prediction skill in summer, it is in fact the only model among the
341 EUROSIP multi-model that has high skill during winter over central Europe
342 and Southern Scandinavia for seasonal temperature and over the British Isles
343 and south of Scandinavia (close to North and Baltic seas) for seasonal precipi-
344 tation. Thus, if a strategy to choose the best model for each region is adopted,
345 MF would add value to the overall EUROSIP multi-model as the preferred
346 model for these regions for the winter season.

[Fig. 3 about here.]

For seasonal temperature during summer, it can be seen that different models are preferred over different regions of Europe. Glosea5 is the preferred model for seasonal temperature over the Scandinavian region, ECMWF over the Mediterranean region and NCEP over the British Isles and East-Central Europe. The skill for seasonal precipitation forecast is mostly noisy and scattered among the four models. The superior correlation of MF for winter over Western Europe, the British Isles and parts of Scandinavia can be noticed.

3.2 Assessment of Probabilistic Ensemble Forecasts

3.2.1 Ensemble Performance of Individual Models

While the ensemble mean in general is the best available estimate of future conditions, the Probabilistic Ensemble Forecasts (PEFs) can provide further information about the distribution of the potential outcome of a prediction. This can be verified using a skill score based on the CRPS, which assesses the relative improvement of the PEFs over climatology to reliably and accurately predict differing observations (Gneiting 2011). In this section, the PEFs drawn from each individual EUROSIP model are assessed using the CRPSS. Figure 4 shows the prediction skill of individual EUROSIP models based on CRPSS for seasonal temperature. During winter, ECMWF exhibits low but significant skill over the British Isles. Other than that, the skill is very limited for all models during winter. For summer, the models exhibit CRPS skill over similar regions as where they exhibited ACC skill, although the skill based on CRPS is lower due to its stringent scoring rule. Glosea5 exhibits skill over Northern Scandinavia and Southern Europe. ECMWF exhibits skill over the Mediterranean regions and South-Eastern Europe. NCEP exhibits very limited significant skill over the East-Central Europe and MF does not exhibit statistically significant CRPS skill over Europe.

We noted earlier that CRPS is known to inflate skill for models with larger ensemble size and ECMWF with the largest ensemble size indeed exhibits the highest skill based on CRPS over most of Europe for seasonal temperature. To verify whether this high skill of ECMWF is due to its larger ensemble size, we calculated the proposed FCRPSS (Ferro 2014; Fricker et al. 2013) for each individual models and show similar results in Figure 5. Additionally, we calculated CRPSS for all individual models using first, only 9 members, and then 15 members (See Fig. S3-4 in supplementary section). Based on these results, we note two things - (1) even with the reduced ensemble size, ECMWF still exhibits higher significant skill and (2) when accounting for the ensemble size, the skill of all CFSs remains over the same regions but is lowered. Thus, the higher skill of ECMWF cannot be attributed solely to its larger ensemble members. This is also true for Glosea5 and NCEP with ensemble size greater than that of MF as well as for MF, when comparing between its 9 and 15 ensemble members. Hence, these models are accurately and reliably capturing

389 atmospheric dynamics to the extent that they perform better than climatology
390 over the parts of Europe where they exhibit skill.

391 [Fig. 4 about here.]

392 [Fig. 5 about here.]

393 Figure 6 shows the results of the same evaluation as that of Figure 4 but for
394 seasonal precipitation. It is seen from these maps that none of the individual
395 CFSs accurately predict seasonal precipitation over Europe for both seasons.
396 Additionally, no noticeable change in skill is found in terms of FCRPSS (see
397 Figure 7) as well as CRPSS with just 9 and 15 ensemble members (see Fig.
398 S5-6 in supplementary section). Thus, PEFs of EUROSIP for seasonal precip-
399 itation show very limited skill over Europe.

400 [Fig. 6 about here.]

401 [Fig. 7 about here.]

402 *3.2.2 Multi-Model Ensemble Performance*

403 The PEFs of the EUROSIP multi-model for seasonal temperature and precip-
404 itation are obtained by taking ensemble anomalies from all four CFSs (118 for
405 winter and 114 for summer). Then, the CRPSS is calculated for the resulting
406 multi-model PEFs to assess whether such multi-model provides higher pre-
407 diction skill than the single best CFS. Based on the results shown in Figure
408 8, a superior predicting skill is not gained by this combination method with
409 respect to the single best model. For both seasons and for both climate vari-
410 ables, the CRPSS of the multi-model PEFs is mostly lower than that of the
411 single best model. For winter seasonal temperature over the British Isles, the
412 CRPSS is much higher for ECMWF than it is for the multi-model PEFs. Simi-
413 larly, significant skill exhibited by PEFs of GloSea5 over Northern Scandinavia
414 for seasonal temperature during summer is not seen in the multi-model PEFs.
415 Some significant skill is gained by multi-model PEFs in parts of Southern Eu-
416 rope and Central Europe during winter and summer, respectively. However,
417 overall the decrease in skill is more evident when compared to that of the
418 single best model. This is seen more clearly in Figure 9, where the maps show
419 only the maximum positive CRPSS among the individual CFSs and the multi-
420 model for seasonal temperature (not shown for seasonal precipitation due to
421 very low skill).

422 [Fig. 8 about here.]

423 [Fig. 9 about here.]

3.3 Assessment of Optimal Method of Combining Forecasts

In this section, weighting each of the individual CFS of EUROSIP with different techniques is tested to determine to which extent the past performance of these CFSs can be utilized to make better predictions of seasonal climate. The maps here evaluate the prediction skill of the three WMM systems (as described in Section 2c) for seasonal temperature (Figure 10) and seasonal precipitation (Figure 11) over Europe. As seen in both figures, a simple MMM is the best combination for SCFs over most of Europe and its predicting skill is significantly higher during summer than winter. Some exceptions can be seen however. For example, BOCM model performs better over the British Isles (most notably over Ireland) during winter. During summer, the skill of BOCM is also higher over parts of Spain, Northern Scandinavia, north of Eastern Europe and over the south of Black sea. The CAWM model shows limited competing skill, although higher skill over the British Isles and along the coastlines of North sea over Western Europe.

When compared to that of the single best models on any given grid point, it is difficult to interpret the skill of these WMMs as it is sensitive to the location and the combination method. Over the entire region of Europe as shown in the map, notably fewer negative skill is exhibited by MMM during winter and by CAWM during summer. However, seasonal temperature skill exhibited during winter over East-Central Russia, the British Isles, Northern coastlines of Russia and Western Europe by Glosea5, ECMWF, NCEP and MF, respectively, is not surpassed by any of the WMMs. During summer, CAWM model exhibits higher skill over the Northern coastlines of Western Europe than that of NCEP, although the skill is not significant. Over south of Europe below Black sea, the skill exhibited by BOCM is higher only by a slight margin.

[Fig. 10 about here.]

[Fig. 11 about here.]

Based on these results, the additional post-processing of forecast data to obtain optimally weighted forecasts is hard to justify. Although, it is also important to note the limitation of these WMMs due to the small sample size. The superiority of a simple MMM technique compared to the unequally WMM in predicting seasonal climate could be attributed to the small sample size available, which for linear combinations methods, is usually not enough to attain robust weights. The coefficients in the linear regression models adapt to the unpredictable variability within the available historical training dataset and thus, performs poorly in the independent data when sample size is limited. Similarly, in the CAWM technique, estimated correlation coefficients can exhibit large uncertainty (Bellprat et al. 2017) and thus, stand as highly volatile weights. Besides, when atmospheric internal variability is large, as in the case of Europe, more information may be lost by inappropriate weighting and therefore, equal weighting may be safer to use. Finally, optimally WMM generally

467 perform better than the single best models only when there is too little in-
468 formation provided by other CFSs within the multi-model (Rodrigues et al.,
469 2014a; 2014b). However, in the case of EUROSIP, each CFS exhibit varied
470 prediction skill based on climate variable, season and region.

471 **4 Discussion**

472 Most of the research on the assessment of forecast skill of SCFs by climate
473 models is done globally and lower skill over the extratropics than in the trop-
474 ics is reported. Very few studies focus exclusively on the European region.
475 The H2020 IMPREX European project has offered an opportunity to study
476 the performance of SCFs over Europe focusing on their usability in adaptation
477 to water-related climate hazards. Robust assessment is needed to objectively
478 evaluate whether SCFs are fit for purpose. To this end, this study aims to fill
479 an important gap in the literature by quantifying the skill of SCFs over Eu-
480 rope using the EUROSIP multi-model and its component operational forecast
481 systems – GloSea5, ECMWF, NCEP and MF.

482 The evaluation is done for seasonal temperature and precipitation forecasts
483 over a period of 21 years (1992-2012) using ACC for deterministic forecasts
484 and CRPSS for probabilistic forecast. The assessment is applied to forecast
485 anomalies (with respect to the models own climatology) against that of obser-
486 vations provided by ERAINT and GPCP for temperature and precipitation,
487 respectively. We also constructed one equally and two unequally WMM sys-
488 tems to evaluate the prospects of model weighting in the context of improving
489 SCFs. Based on the results of this study, limited predictive skill of seasonal
490 temperature and very low skill of seasonal precipitation forecast is found over
491 Europe.

492 For seasonal temperature, the forecast skill of competing models differ
493 based on both regions and seasons. The predictability is higher during sum-
494 mer than winter. The higher skill based on ACC during summer for seasonal
495 temperature forecasts can be associated with the warming trend (results not
496 shown) as the data has not been detrended (Doblas-Reyes et al. 2013). The
497 higher skill during summer can also be because interannual variability tends
498 to be weaker for temperature during summer (Doblas-Reyes et al. 2000). The
499 skill of MF over the Western Europe during winter can be attributed to the
500 extreme soil moisture conditions in South-Central Europe (van den Hurk et al.
501 2012). The low predictive skill for European winter was also reported by Wehrli
502 et al. (2017), who assessed the seasonal temperature forecasts of ECMWF back
503 to 1981. This low skill can be associated with blocking events or winter me-
504 teorological perturbations as well as the misrepresentation of teleconnections
505 leading to erroneous NAO-related signals (Doblas-Reyes et al. 2003; Kim et al.
506 2012; Wehrli et al. 2017). Scaife et al.(2014), however, showed NAO forecasting
507 skill in GloSea5 and thus, attributed the low skill over temperature to overdis-
508 persion, suggesting that increasing the ensemble size would increase forecast
509 skill. However, we show that the skill of a multi-model PEF, which combines

ensembles from all four models, did not surpass the skill of the single best model over most of the grid points.

In the case of winter precipitation, MF has some significant correlation skill over the British Isles, parts of Scandinavia and Italy. Some skill over parts of Eastern Europe is found in ECMWF and GloSea5 during winter, and in NCEP and MF during summer. Other than that, seasonal precipitation predictability over Europe is remarkably low. The results of this study are consistent with that of Branković and Palmer (2000), Doblas-Reyes et al. (2000), Graham et al. (2000), Lavers et al. (2009) and Wehrli et al. (2017), all of whom also assessed the forecast skill of seasonal precipitation and showed lower skill over the extratropics than in the tropics, although not negligible. This predictability for seasonal mean precipitation over Europe is attributed to ENSO and local North Atlantic SST forcings. (Doblas-Reyes et al. 2013; Frías et al. 2010; Lloyd-Hughes and Saunders 2002).

Additionally, our results highlighted the benefits of having a multi-model system given that the skill of each individual CFS varied across locations, seasons and climate variables. The difference in skill exhibited by different models also provides context to further examine and understand the climate phenomena that are differently represented by these models. However, the differing number of ensemble sizes among the CFSs complexifies the comparison between the models. In terms of probabilistic skill assessment using CRPS, it was highlighted by Ferro et al. (2014; 2008) that CRPS inflates the score for models with higher number of ensemble members. We tested the effect of ensemble size on superior predictability by models with higher number of members. First, we calculated the recommended FCRPSS that corrects for systematic bias in skill scores induced by the finite ensemble size (Ferro 2014; Ferro et al. 2008; Fricker et al. 2013). Then, we also compared the skill based on CRPS taking equal numbers (9 and 15) of ensemble members for all CFS. The results show that ECMWF still has higher prediction skill over most of Europe compared to other CFSs. Additionally, we show that increasing the ensemble size contributes to higher and more significant prediction skill, not only for ECMWF, but for all CFSs.

Varied levels of skill among the models has also garnered an interest to combine their forecasts based on historical performance of individual CFSs. In this study however, the multi-model PEFs of EUROSIP did not present a considerable improvement in CRPSS when compared to that of the single best model. This could be because the optimal set of models in the ensemble may vary in time for continuous forecasting scenarios as argued by Krikunov and Kovalchuk (2015). They suggested a method to dynamically select ensemble members in multi-model and demonstrated slight improvement of forecast skill. Further improvement opportunities also lie in the use of different machine learning techniques and artificial neural networks for such dynamic selection procedures.

In case of deterministic multi-models, the two unequally weighted methods, BOCM and CAWM, did not always outperform the simple equally weighted MMM. When compared to the single best model, the skill of WMMs is sensitive

556 to the location and combination method and thus difficult to interpret. The
557 limited sample size in this study can be attributed to this sensitivity. A sample
558 size of 21 years is usually not enough to attain robust weights with linear
559 regression or correlation methods. In principle, optimal weighting can optimize
560 the predictive skill (DelSole et al. 2013). However, the ideal way to combine
561 these forecasts is still far from being trivial. These results are also consistent
562 with Rodrigues et al.(2014), who also found that equally or unequally WMM
563 approaches does not always perform better, especially when a models with
564 lower skill are present in a multi-model system.

565 In contrast to this, Wanders and Wood (2016) showed considerable im-
566 provements through unequally WMM. This can be attributed to their assess-
567 ment of SCFs skill for longer time-period of 30 years (and over large-scale
568 area averages), which yields temporally and spatially stable relationships and
569 allows more robust estimation of weights. Thus, larger homogenous dataset
570 of competing individuals CFSs is critical for the performance of WMMs. The
571 Copernicus Climate Change Services (C3S) by the Copernicus programme
572 comprise all year round seasonal climate forecasts from several state-of-the-
573 art seasonal prediction systems for longer a period. This offers an opportunity
574 to extended the analysis of this study in the near future.

575 All verification scores in the study are obtained from forecast anomalies
576 to filter out the systematic biases existent in climate models that can lead to
577 underestimation of true predictability. However, anomaly correction only par-
578 tially removes these biases. The systematic errors in the model variability can
579 be possibly accounted for by implementing other bias correction and calibra-
580 tion methods (Bazile et al. 2017; Crochemore et al. 2016; Torralba et al. 2017).
581 We implemented a Simple Univariate Bias Correction (SUBC) method, which
582 resulted in further decrease in the prediction skill (See Fig. S7-8 in supplemen-
583 tary section, not shown for precipitation due to the inappropriate Gaussian
584 assumption). This is because SUBC directly adjusts the distribution of fore-
585 casts from CFSs against the observations to match their statistical properties
586 (mean and standard deviation), which requires parameter estimation and fur-
587 ther adds uncertainty in the forecast datasets (Bazile et al. 2017; Manzanas
588 et al. 2017). Besides, the choice of bias correction method also depends on the
589 climate variable and the intended use of its seasonal forecast. In addition, cal-
590 ibration of pattern errors and probability forecast adjustments generally also
591 require longer training datasets.

592 **5 Conclusion**

593 This study evaluated forecasts from the EUROSIP multi-model database and
594 highlighted the limited skill of forecasting seasonal mean temperature and pre-
595 cipitation over Europe for winter and summer. Based on these results, for a
596 comprehensive assessment of the prediction skill of seasonal climate variables,
597 it is recommended to analyze the prediction skill of both individual and com-
598 bined multi-model systems to identify the one with the highest skill over any

599 given area and/or season. The overall lack of forecast skill is because most
600 of the mechanisms driving changes in seasonal temperature and precipitation
601 over Europe are a topic of active research still. In addition, even if the mech-
602 anism are well known such as the NAO, they are challenging to represent by
603 coupled global climate models. Thus, significant improvements in the opera-
604 tional climate forecast systems are required through an improvement of the
605 initial conditions and the realism of the climate models. Improved predictabil-
606 ity can benefit decision making based on these forecasts and better prepare
607 the European society against the anticipated water-related climate hazards.

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Correlation of Seasonal Temperature for Individual EUROSIP Models

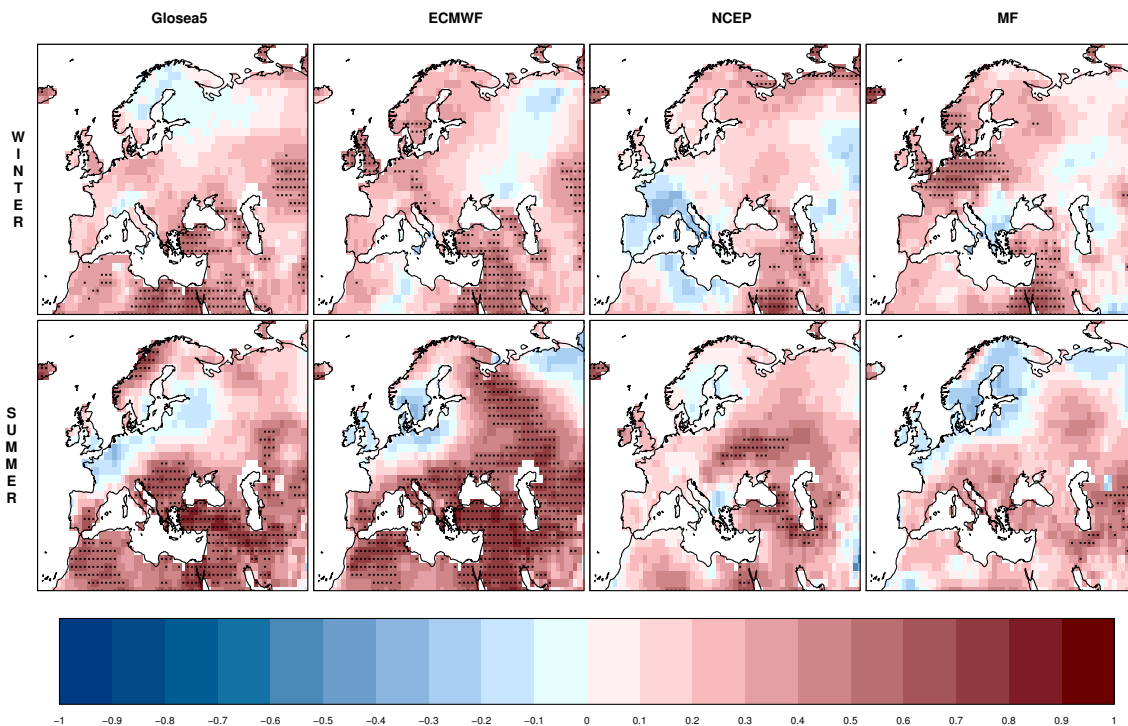


Fig. 1: Anomaly Correlation Coefficient (ACC) between the predicted ensemble mean of each individual climate model of EUROSIP and the observed seasonal winter (DJF; top row) and summer (JJA; bottom row) temperature obtained from ERAINT over the European region (20° W-70° E and 25° N-75° N) for the period 1992-2012. The individual climate models are Glosea5, ECMWF, NCEP and MF (from left to right; see details in section 2). Forecasts are initialized in November for DJF and in May for JJA. Areas covered in red are indicative of positive correlation, while areas covered in blue indicate negative correlation. Dots in each grid point indicate significant positive correlation at 5% significance level using one-sided Student t-test and controlling for false discovery rate.

Correlation of Seasonal Precipitation for Individual EUROSIP Models

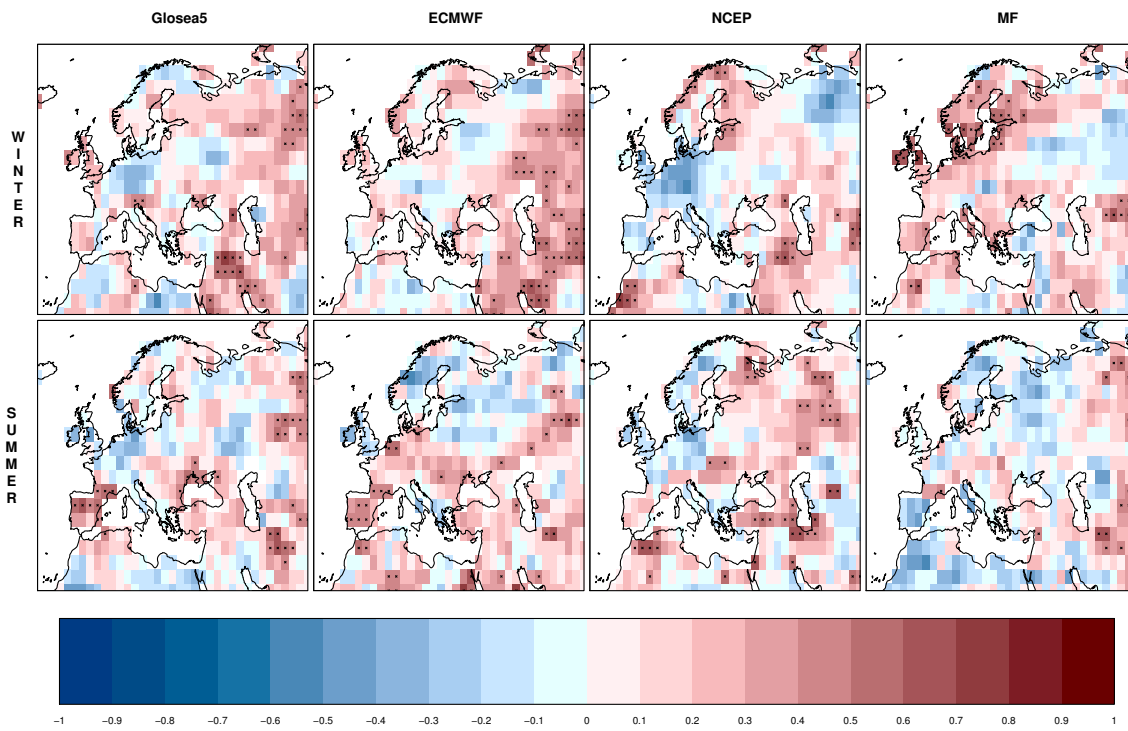


Fig. 2: Same as Fig.1 but for precipitation and reference data obtained from GPCP

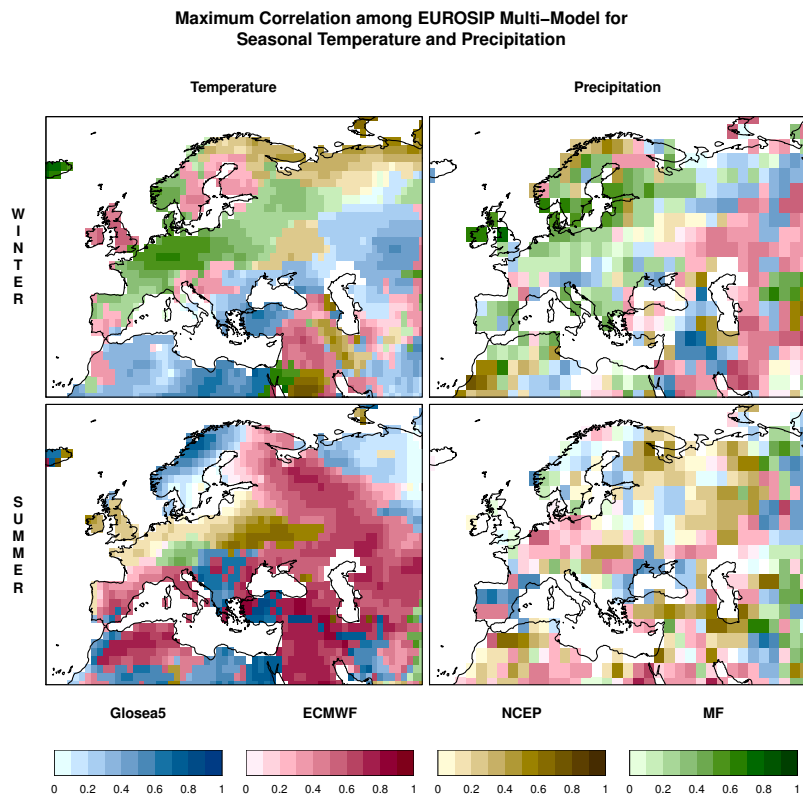


Fig. 3: Maximum positive Anomaly Correlation Coefficient (ACC) among the four individual models from EUROSIP. ACC for each model is calculated between their respective predicted ensemble mean anomalies and the anomalies of the observed temperature obtained from ERAINT (left) and of precipitation obtained from GPCP (right) for winter (DJF; top row) and summer (JJA; bottom row) seasons over the period 1992-2012. Blue, red, yellow and green colors indicate that the maximum correlation is obtained for GloSea5, ECMWF, NCEP and MF respectively. Negative or 0 correlations appear in white.

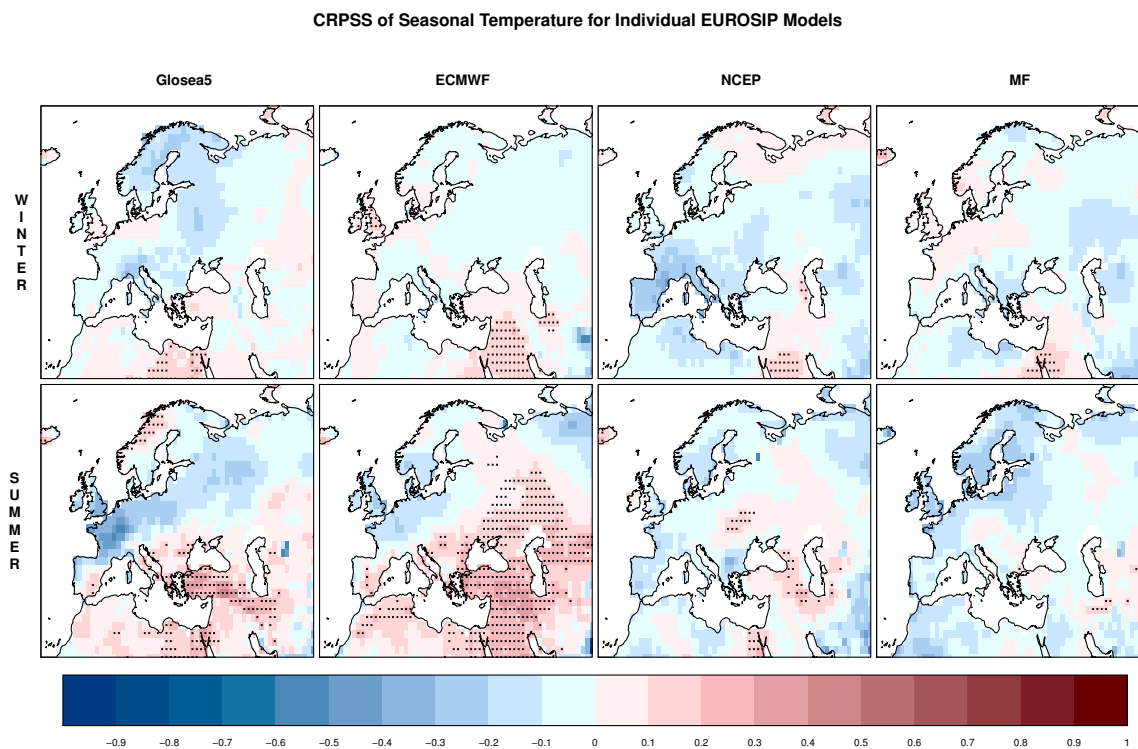


Fig. 4: CRPSS of the probability ensemble forecasts of each individual climate model of EUROSIP with climatology used as reference forecast obtained from ERA-Interim for winter (DJF; top row) and summer (JJA; bottom row) seasonal temperature over the European region (20° W- 70° E and 25° N- 75° N) for the period 1992-2012. The individual climate models are GloSea5, ECMWF, NCEP MF (from left to right). Forecasts are initialized in November for DJF and in May for JJA. Areas covered in red are indicative of positive CRPSS, suggesting skill better than climatology. Areas covered in blue indicate worse skill than climatology. Dots in each grid point indicate significant positive CRPSS using the standard deviation of the skill score, approximated by propagation of uncertainty at 95% confidence interval.

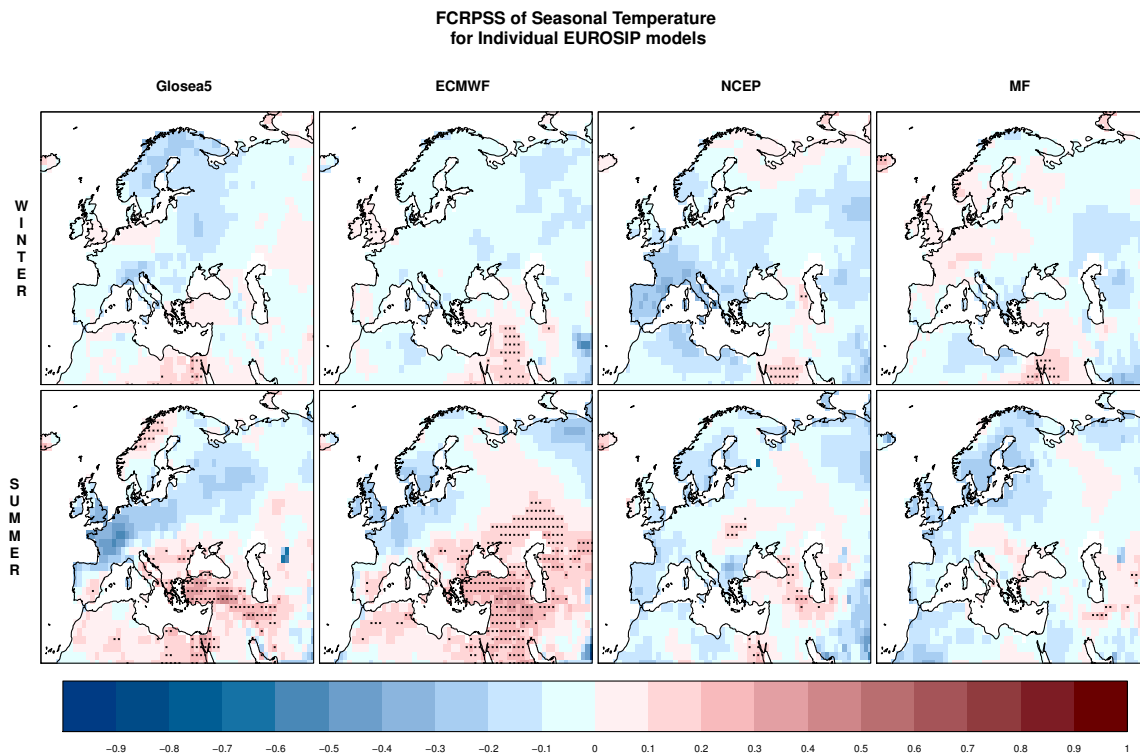


Fig. 5: FCRPSS of the probability ensemble forecasts of each individual climate model of EUROSIP with climatology obtained from ERAINT used as reference forecast for winter (DJF; top row) and summer (JJA; bottom row) seasonal temperature over the European region (20° W-70° E and 25° N-75° N) for the period 1992-2012. The individual climate models are Glosea5, ECMWF, NCEP and MF (from left to right). Forecasts are initialized in November for DJF and in May for JJA. Areas covered in red are indicative of positive CRPSS, suggesting skill better than climatology. Areas covered in blue indicate worse skill than climatology. Dots in each grid point indicate significant positive CRPSS using the standard deviation of the skill score, approximated by propagation of uncertainty at 95% confidence interval.

CRPSS of Seasonal Precipitation for Individual EUROSIP Models

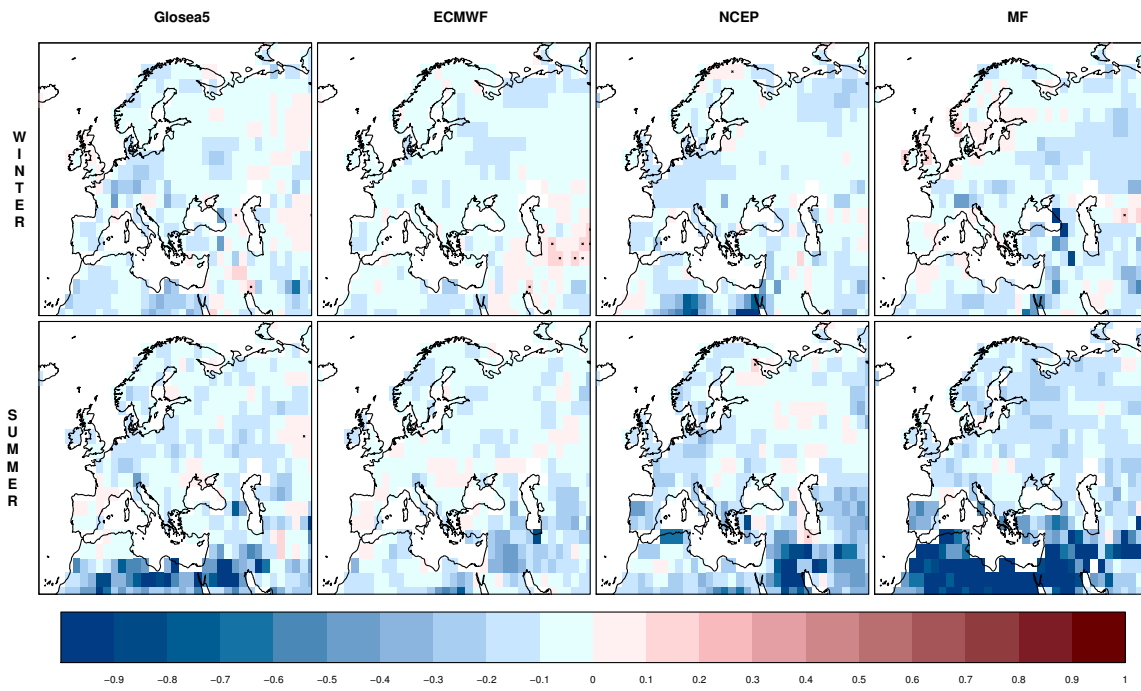


Fig. 6: Same as Fig.4 but for precipitation and reference data obtained from GPCP.

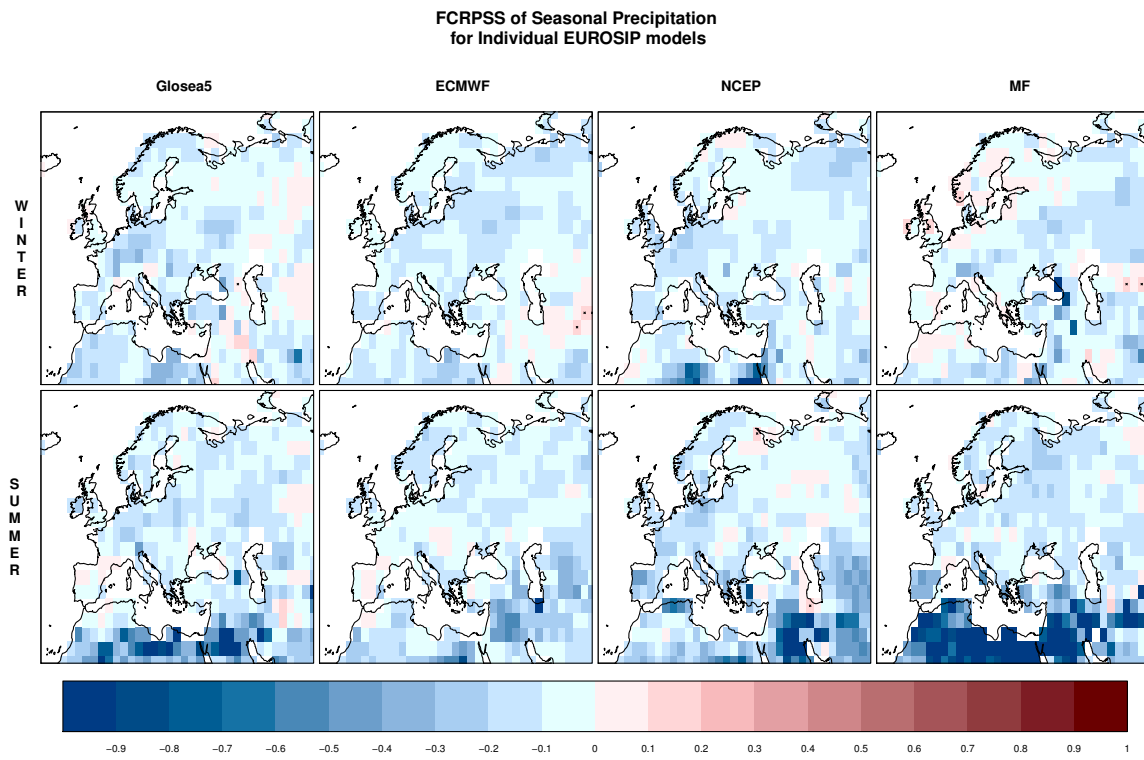


Fig. 7: Same as Fig. 5 but for precipitation and reference forecast obtained from GPCP.

CRPSS of Seasonal Temperature and Precipitation for EUROSIP Multi-Model

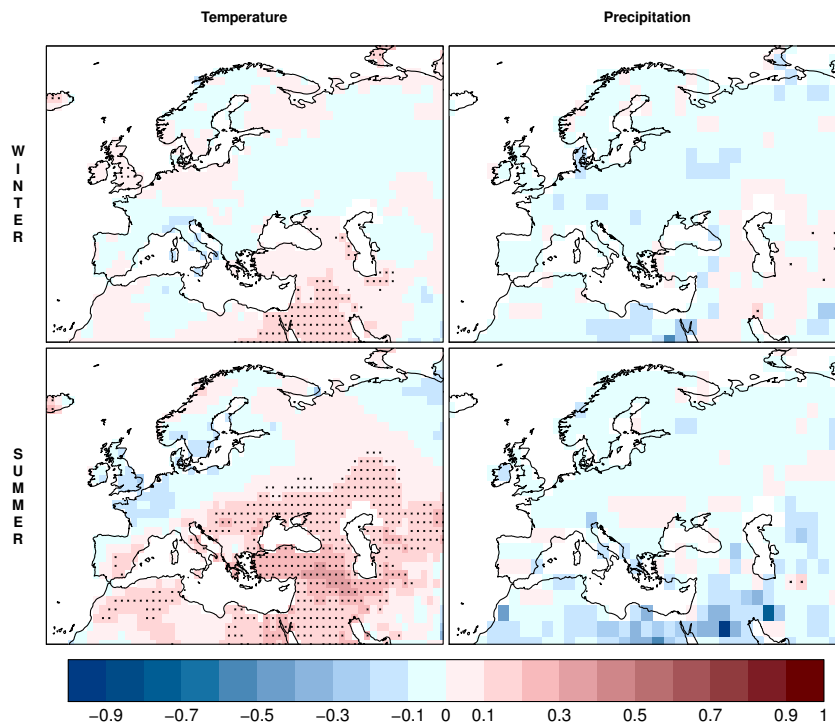


Fig. 8: CRPSS of the probability ensemble forecasts from all four individual climate models of EUROSIP (GloSea5, ECMWF, NCEP and MF) treated as one single model with climatology used as reference forecast obtained from ERA-Interim for temperature (left) and GPCP for precipitation (right) for winter (DJF; top row) and summer (JJA; bottom row) seasons over the European region (20° W-70° E and 25° N-75° N) for the period 1992-2012. Forecasts are initialized in November for DJF and in May for JJA. Areas covered in red are indicative of positive correlation suggesting skill better than climatology. Areas covered in blue indicate worse skill than climatology. Dots in each grid point indicate significant positive CRPSS using the standard deviation of the skill score, approximated by propagation of uncertainty at 95% confidence interval.

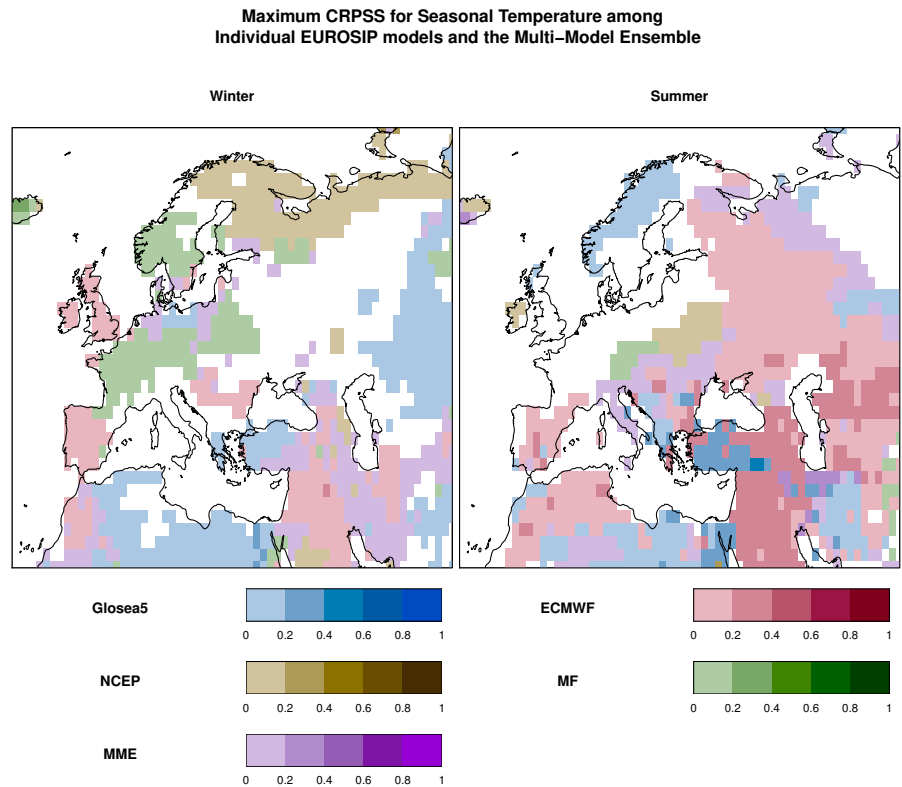


Fig. 9: Maximum positive CRPSS among the four individual models from EUROSIP and the multi-model with climatology used as reference forecast obtained from ERAINT for winter (DJF; left) and summer (JJA; summer) over the European region (20° W-70° E and 25° N-75° N) for the period 1992-2012. Forecasts are initialized in November for DJF and in May for JJA. Blue, red, yellow, green and purple colors indicate that the maximum CRPSS is obtained for GloSea5, ECMWF, NCEP, MF and the multi-model, respectively. Negative or 0 correlations appear in white.

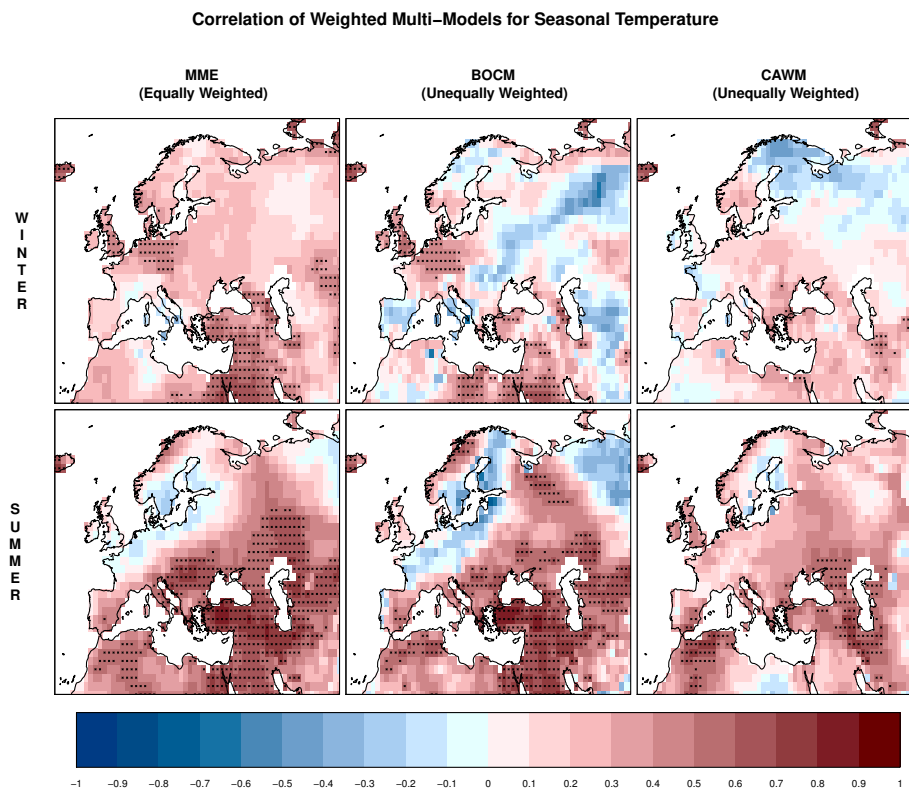


Fig. 10: Anomaly Correlation Coefficient (ACC) between the predictions obtained from the three Weighted Multi-Model (WMM) systems of EUROSIP - Multi-Model Mean (MMM), Best OLS Combination Model (BOCM) and Correlation As Weights Model (CAWM; from left to right) and the observed seasonal winter (DJF; top row) and summer (JJA; bottom row) temperature obtained from ERAINT, respectively, over the European region (20° W- 70° E and 25° N- 75° N) for the period 1992-2012. Areas covered in red are indicative of positive correlation, while areas covered in blue indicate negative correlation. Dots in each grid point indicate significant positive correlation at 5% significance level using one-sided Students *t*-test. Details on the construction of each WMM system are given in Section 2c.

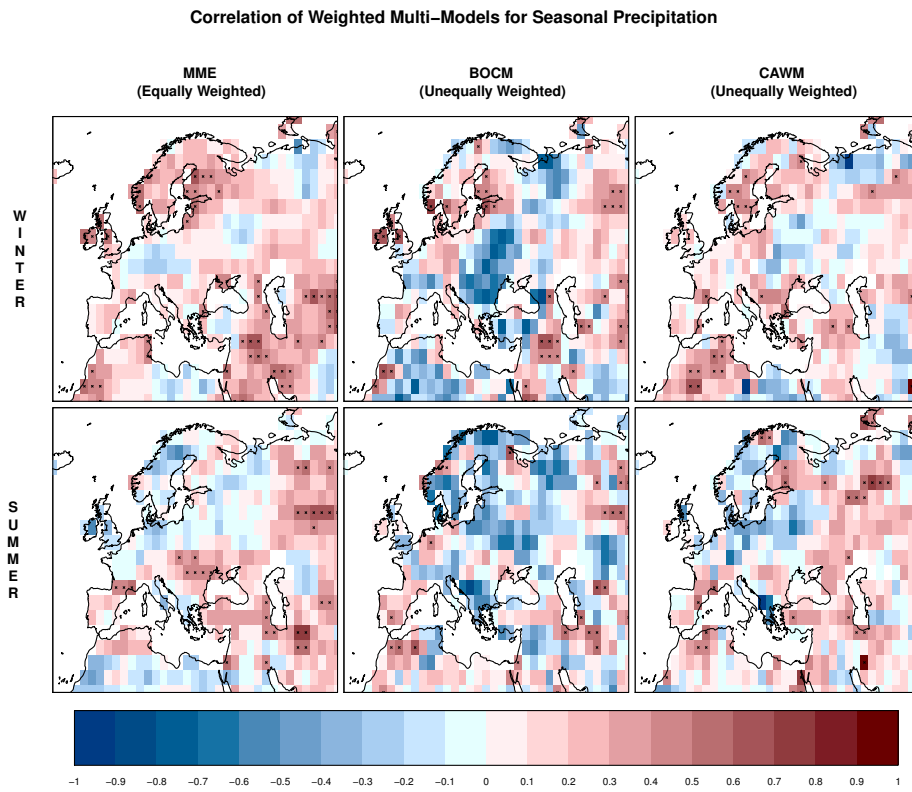


Fig. 11: Same as Fig.10 but for precipitation and reference data obtained from GPCP.

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Table 1: Individual climate models of EUROSIP multi-model seasonal forecasting system

Climate Model	No. of Ensemble Members	Resolution (in Gaussian grid)
GloSea5	24	512x256
ECMWF	51	432x325
Meteo France	15	256x128 for temperature
		360x181 for precipitation
NCEP	28 for winter	384x190
	24 for summer	
Reference Dataset		Resolution (in Gaussian grid)
ERA-Interim for temperature		512x256
GPCP for precipitation		144x72