

AMERICAN METEOROLOGICAL SOCIETY

Journal of Atmospheric and Oceanic Technology

EARLY ONLINE RELEASE

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The DOI for this manuscript is doi: 10.1175/JTECH-D-13-00083.1

The final published version of this manuscript will replace the preliminary version at the above DOI once it is available.

If you would like to cite this EOR in a separate work, please use the following full citation:

Fietzek, P., B. Fiedler, T. Steinhoff, and A. Körtzinger, 2013: In situ quality assessment of a novel underwater pCO2 sensor based on membrane equilibration and NDIR spectrometry. J. Atmos. Oceanic Technol. doi:10.1175/JTECH-D-13-00083.1, in press.

In situ quality assessment of a novel underwater pCO_2 sensor based on membrane equilibration and NDIR spectrometry

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Abstract

We present a detailed quality assessment of a novel underwater sensor for the measurement of $\rm CO_2$ partial pressure ($p\rm CO_2$) based on surface water field deployments carried out between 2008 and 2011. The commercially available sensor, which is based on membrane equilibration and NDIR spectrometry is small and can be integrated into mobile platforms. It is calibrated in water against a proven flow-through $p\rm CO_2$ instrument within a custom-built calibration setup. The aspect of highest concern with respect to achievable data quality of the sensor is the compensation for signal drift inevitably connected to absorption measurements. We use three means to correct for drift effects: (i) a filter correlation or dual-beam setup, (ii) regular zero gas measurements realized automatically within the sensor and (iii) a zero-based transformation of two sensor calibrations flanking the time of sensor deployment.

Three sensors were tested against an underway $p\rm CO_2$ system during two major research cruises providing an in situ temperature range from 7.4 to 30.1°C and $p\rm CO_2$

values between 289 and 445 µatm. The average difference between sensor and reference

 pCO_2 was found to be -0.6 \pm 3 µatm with a RMSE of 3.7 µatm.

1. Introduction

The measurement of dissolved carbon dioxide (CO ₂) in seawater is important and
valuable for a large number of scientific, industrial as well as socio-economic issues.
Major scientific interest is related to the anthropogenic increase of atmospheric CO_2
concentrations and the resulting oceanic uptake of this most important anthropogenic
greenhouse gas (Sabine et al. 2004; Rogner et al. 2007). The exchange of CO ₂ across the
air-sea interface and the dynamics and trends of the carbon cycle in coastal and open
ocean regions (Takahashi et al. 2009; Gruber et al. 2010) as well as in the interior ocean
are key aspects of current marine carbon cycle research.
Dissolved CO ₂ reacts with water to form carbonic acid (H ₂ CO ₃), which rapidly
dissociates into hydrogen (H ⁺), bicarbonate (HCO ₃ ⁻), and carbonate ions (CO ₃ ²⁻). The
exact speciation within the marine CO ₂ system, i.e. between the above species of the
carbonic acid dissociation system, strongly affects the pH of seawater (Millero 2007) and
is of major interest. An increasing amount of dissolved CO2 therefore leads to a
decreasing pH, a process also referred to as ocean acidification in the context of the
anthropogenic CO ₂ transient. Its impact on calcifying organisms as well as on the
physiology and reproduction of other marine species is presently not well understood
(Doney et al. 2009). The potential long-term influence of acidification on fisheries but
also carbon capture and underwater storage scenarios (IPCC 2005) show the socio-
economic relevance of CO ₂ measurements and highlight the demand for baseline
monitoring of CO ₂ parameters in the ocean.
Since the thermodynamic constants of the marine carbonate systems are known
rather precisely (e.g., Millero 2007), the marine CO ₂ system can be fully determined by

measurement of any two of the following variables: dissolved inorganic carbon (DIC), total alkalinity (TA), pH and CO_2 fugacity (fCO_2) or CO_2 partial pressure (pCO_2 ; Millero 2007). Although the sole measurement of pCO_2 is not sufficient to fully characterize the marine CO_2 system, pCO_2 is still a useful parameter on its own: It is the determining factor for air-sea gas exchange and responds sensitively to biogeochemical processes such as photosynthesis and respiration. Hence pCO_2 is both, a valuable stand-alone measurement parameter and a useful measured variable within multi-parameter measurements for the determination of the CO_2 system. So far only pCO_2 and pH can be measured autonomously with commercial underwater sensors (DeGrandpre et al. 1995; Seidel et al. 2008; Martz et al. 2010). Sensors for autonomous measurements of the other carbonate system parameters TA and DIC (Byrne et al. 2002; Wang et al. 2007) and even for direct measurements of carbonate ions are under development (Byrne and Yao 2008). Due to a lack of a commercially available underwater pCO_2 sensor with a sufficient accuracy and platform integratability expressed by adequate dimensions, an appropriately short response time at surface and at depth as well as by the ability for continuous measurements, we have advanced the development of an autonomous and commercially produced underwater pCO₂ sensor (HydroC-CO₂, CONTROS GmbH, Kiel, Germany). Its versatility and specifications allow for integration into various platforms and provide a suitable tool for pCO_2 measurements at an improved spatial and temporal resolution in the water column. Here, we present a comprehensive overview of the sensor's measurement principle, its key components and its calibration. We also assess the achievable data quality by means of a detailed analysis of data from surface water field applications. Measurements against a proven flow-through system only represent

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one application of the sensor, but at the same time enable the clearest assessment of the sensors data quality. Many further applications in the water column are possible and best require a dedicated discussions such as the measurements on a profiling float presented in Fiedler et al. (2013).

a. Suitable sensor platforms

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Beside classical mobile sensor platforms on the one hand, such as voluntary observing ships (VOS) or research vessels, and stationary buoys and moorings on the other hand, innovative mobile platforms are receiving growing interest (Fietzek et al. 2011). These platforms, e.g. autonomous underwater vehicles (AUV), profiling floats and gliders, provide an increased autonomy, mobility and versatility as opposed to the classical carriers. They allow for a more cost-efficient data collection on so far largely unexplored temporal and spatial scales. These newly accessible scales are of high interest for the investigation of various biogeochemical processes, making these modern mobile platforms desirable carriers for pCO_2 sensors (Gruber et al. 2010). By equipping autonomous platforms as a regional or global sensor array with high spatiotemporal resolution, such as the prominent Argo project, a high scientific potential can be achieved (Roemmich et al. 2009; Johnson et al. 2009; Fiedler et al. 2013). However, in order to be easily installed on modern mobile platforms, new sensors need to cope with the demands of the more complex carriers and fulfill the more stringent requirements with respect to payload capacity, power availability, response time, etc. A more detailed discussion of these facts and circumstances can be found in Fietzek et al. (2011).

b. pCO_2 measurements

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1950's. These were based on wet-chemical pH-determination behind a Teflon™ membrane (Stow et al. 1957; Severinghaus and Bradley 1958). Today optical CO₂ measuring techniques are prevailing. One method is to detect and quantify CO₂ molecules within an equilibrated gas stream by means of direct absorption in the infra red (IR) region of the electromagnetic spectrum. Another technique is an indirect measurement making use of the pH affecting property of CO₂ by applying spectrophotometry within an equilibrated pH-sensitive dye solution of known characteristics (DeGrandpre et al. 1995; Lefèvre et al. 1993). An overview of current sensor techniques for carbonate system species can be found in Byrne et al. (2010). A technical evaluation of pCO_2 sensors that also includes two sensors of the type discussed here is presented in Tamburri et al. (2011). Between the evaluation and the measurements discussed herein the status of the sensors was mainly improved through an optimized calibration process and data processing. The usage of underway flow-through instruments to measure pCO_2 both in the oceanic surface layer and in the atmosphere dates back to the 1960s (Takahashi 1961). While in the beginning the application of these systems was restricted to research vessels, current, improved systems are suitable also for application on unattended platforms such as VOS (Watson et al. 2009; Pierrot et al. 2009; Pfeil et al. 2013). Continuous optimization of the overall setup and the components used lead to some standardized

design that is nowadays also commercially available. The key components of such a

flow-through instrument are the air-seawater equilibrator and the IR gas analyzer. The

The first measurements of dissolved CO₂ were developed for medical applications in the

equilibrator is used to achieve partial pressure equilibrium between the dissolved gases in a continuous stream of sea water and air that is re-circulated between the equilibrator and a benchtop IR analyzer. The achievable measuring accuracy is 2 μ atm for surface seawater pCO_2 measurements (Pierrot et al. 2009). The flow-through system data are commonly reported as pCO_2 although the fCO_2 is the value suggested for most accurate carbonate system calculations. Based on approximated expressions the fCO_2 , which considers the slightly non-ideal behavior of CO_2 in the gas phase, can be calculated from temperature and pCO_2 , which presumes ideal gas behavior.

2. Sensor principle and description

a. Development and design

The development aims of the pCO_2 sensor among others were to obtain (i) a versatile and autonomous sensor that could be deployed on a profiling float with (ii) a response time of less than 5 min, (iii) an accuracy better 5 μ atm, and (iv) a stability and reliability that would allow for long-term deployments of several months.

Since the developed sensor is based on the same measuring principle as proven flow-through systems, it has the same key components: an equilibrator and an IR CO₂ detector. A planar, semi-permeable membrane with a silicone active layer is installed in the head of the sensor. It acts as an equilibrator as well as a phase separator between the ambient water and an internal headspace. The sensor is commonly equipped with a water pump that provides a continuous seawater flow to the membrane and thus reduces the thickness of the static boundary layer in front of the membrane. By that, the response time is effectively shortened and made independent of a relative movement between the

membrane and the surrounding water. In order to withstand high hydrostatic pressures the membrane is mechanically supported from behind with a sintered metal disc. A gas pump continuously circulates air between the membrane equilibrator and a non-dispersive IR detector (NDIR). Figure 1 shows a model of the sensor and provides a schematic overview of its setup. The gas tightness of the internal gas stream as well as of the integrated valves is checked thoroughly prior to calibration. Opposed to pCO₂ underway systems in which gas stream leakages are a major source for measuring errors (Pierrot et al. 2009), the biggest "leak" within the gas stream of this sensor remains the equilibration membrane; the high gas permeability of the membrane related to the volume of the gas stream compensates for possible influences caused by small leaks. If bigger leakage occurs within the gas stream, the entire, much larger internal gas volume of the sensor will be equilibrated, leading to noticeably slower response times but not necessarily biases. The gas circuit also features a specially developed gas heater upstream of the NDIR detector whose heating control system is also used to stabilize the temperature of the IR detector. The gas heater buffers seawater temperature gradients ($\Delta T_{\text{in-situ}}$) in such a way that large $\Delta T_{\text{in-situ}}$ are damped to a much smaller gas temperature gradient. We choose to set the control temperature just high enough for the heating control circuit to keep the controlled temperature stable even at the maximum in situ temperature expected during deployment. When the control temperature is set higher than necessary, avoidably high power consumption is the consequence and the abundant absolute temperature differences between the surrounding water and the internal gas becomes larger. In addition to the gas heater and the temperature stabilization, the sensor's pressure housing is thermally insulated and temperature sensitive components are separately protected.

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Within the gas stream sensors for pressure, temperature and relative humidity (RH) are installed to determine the conditions within the NDIR detector as well as behind the membrane. Their exact position was chosen upon laboratory tests to be most beneficial for their consideration within the IR sensor data analysis. The quality and a deep understanding of the NDIR detector are crucial for the data quality of both, underway instruments as well as the new underwater sensor. All additional components within the underway instrument's and the underwater sensor's gas circuit beside the equilibrator and NDIR detector, such as the additional sensors mentioned above, are required for accurate and precise absorption measurements and allow for preferably long deployments.

In contrast to common practice of underway instruments the absorption measurement within the sensor is carried out in wet air and without interrupting the gas flow for measurement. In the underwater sensor regular zero gas measurements can be carried out. Therefore valves are included into the circuit that lead the pumped air through a soda lime cartridge instead of the membrane equilibrator at desired intervals (see Fig. 1). In the presence of water vapor soda lime scrubs the CO₂ binding it as calcium carbonate (CaCO₃) thus creating a zero gas with respect to CO₂.

The sensor operates by consecutively switching through different intervals, the durations of which can be individually set. As soon as the sensor is powered it starts with a warm-up interval followed by continuous repetition of measuring cycles. One measuring cycle consists of three intervals: zero, flush and measure. The warm-up interval is only passed through once after the sensor has been turned on. The required warm-up time depends on the water temperature and the supplied voltage (c.f. Table 1).

During the warm-up, the water pump is disabled and data are neither transmitted nor stored in the internal logger. During zero intervals, a zero CO₂ gas is created as described above and the sensor provides the current zero reading used later for drift correction. The zeroing typically lasts a few minutes and repetition is recommended at least every 12 h. Data stored during that time are flagged for easy handling during analysis. Flush intervals are only used to flag data acquired during the signal recovery from the zero value to the ambient pCO_2 reading. Technically the measuring process of the sensor does not differ between the flush and the subsequent measuring interval. The time the sensor needs for full equilibration depends on the sensor configuration and the environmental conditions, mainly the water temperature. The warmer the water, that faster the response time and thus the shorter the required flush time. Laboratory and field tests for the response time determination of the sensor as configured within the deployments presented here indicate a linear dependency of response time on water temperature at a slope of the order of -1 s per 1 °C (data not shown here). Sensor response to a step input can be well described by first order kinetics and a corresponding exponential fit. Any response time given here hence represents a time constant or a $t_{63\%}$ depending on which formulation is favored. Isothermal pressure vessel experiments up to 200 bar have neither shown indications for a pressure hysteresis, nor could a significant pressure influence on the response time be identified (data not shown here). The actual response time of the sensor can be derived based on the course of the signal recovery during the flush interval as applied by Fiedler et al. (2013). It is typically faster than the 2 min response time of the flow-through system (Pierrot et al. 2009) and allows for measurements on moving platforms; especially when a response time correction is applied to derive the "true" ambient pCO_2 from the time-

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lagged sensor signal (Fiedler et al. 2013). A more detailed analysis of the sensor's response characteristics is in preparation. During flush and measurement intervals, the water pump is active and data are recorded as configured. The interval settings simplify the data processing and provide the means to generate measurement data from a fully equilibrated and internally temperature stabilized sensor. Figure 2 shows the sensor signal during calibration with the zero, flush and measurement intervals indicated.

The comparatively fast response time, small size and operability of the sensor allow for deployments on various platforms. Due to its design and compared to classical flow-through systems the maintenance intensity and the risk for leaks in the gas stream are low. If used in situ, the sensor lacks the demand for an accurate water temperature probe as it is crucial for typical flow-through systems (Körtzinger et al. 2000). In order to derive the actual amount of dissolved CO₂ from the measured partial pressure, the sensor is commonly deployed together with a CTD probe, as the solubility of CO₂ depends on temperature and salinity (Weiss 1974).

b. Specifications

Table 1 lists the specifications of the developed sensor as applicable for the measurements discussed in this work. Different data communication options are feasible. An internal data logger is optional that can either be used as a stand-alone memory (e.g. Saderne et al. 2013) or for backup storage (e.g. Fiedler et al. 2013). A sleep mode function further facilitates autonomous installations. The sensor development also comprised a surface water flow-through version of the sensor which is not further addressed here.

c. CO_2 measurements by means of NDIR spectrometry

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The properties of the NDIR detector are relevant for the overall performance of the sensor. Drift, cross-sensitivities and the signal-to-noise ratio (S/N) of the NDIR detector directly affect the data quality.

NDIR spectrometry in general is a proven direct measuring technique for mole fractions of gases absorbing in the IR. It is non-destructive and traceable to standards. NDIR detectors for CO₂ have small dimensions (several cm) and moderate power consumption (here: about 0.5 W) that allow for easy integration. In addition they provide good mechanical strength and are unaffected by vibrations if realized without moving parts. Due to their high selectivity and limited cross sensitivity NDIR detectors are well suited for qualitative analysis. NDIR detectors are composed of three main components: light source, absorption/beam path and detector. Each of them has different influences on the final sensor signal and depends differently on environmental variables, e.g. temperature. This may lead to complex overall sensor properties. Their choice defines the S/N, the sensitivity and the measurement range (detection limit and upper range value) of an NDIR unit. The basic idea is to get enough light energy at the desired wavelength to the detector and to make the absorption path sufficiently long such that changes in absorbed light intensity can be clearly resolved by the detector and amplifying electronics. Various options exist to optimize and dovetail these components.

The linear relation between the transmitted light intensity, I, and molecule concentration, c, as described by the Beer–Lambert Law makes the technique suitable for quantitative analysis:

$$I = I_0 10^{-\varepsilon cl} , \qquad (1)$$

with I_0 being the initial light intensity, l the distance the light travels through the absorbing medium and ε the molar absorption or extinction coefficient of the target molecule to be detected. According to (1) the extinction or absorbance, A, is defined to be directly proportional to the molecule concentration in the medium:

$$A = \log\left(\frac{I_0}{I}\right) = \varepsilon cl. \tag{2}$$

In reality there is a small non-linearity between the absorbance measured by the NDIR detector and molecule concentration due to the fact that the Beer–Lambert law is only defined for a single wavelength of infinitesimal small width, while in applications spectral dependencies occur (Wiegleb et al. 2001).

The NDIR unit used features a black body radiator as a broad band IR light source and a pyro-electrical IR detector which requires pulsed operation of the IR emitter. Interference filters in front of the detectors select the desired measurement as well as reference wavelength and together with electronics for control and evaluation complete the single beam dual wavelength NDIR detector. CO_2 is typically sensed around 4.26 μ m where it shows its highest absorption due to its fundamental asymmetric stretching vibration (2349.1 cm⁻¹) and the reference wavelength is chosen to be placed in the water vapor window at around 4.0 μ m.

Within an NDIR detector, several temperature influences exist that either require a compensation, calibration or stabilization. The spectral properties of the filters are temperature dependent: The central filter wavelength can shift as well as the transmission width can vary. A temperature influence on the sensor can occur in the form of thermal noise, a thermal background signal and changes in sensitivity. In the case of a dual beam setup the influences might even be different for both channels. The emission properties

and emitted intensities of the light source show a temperature influence as well. Thermal expansion of the cuvette or other mechanical deformations of the absorption path may have an additional effect on the measurements. Finally temperature dependencies of the analog electrical components directly behind the detector need to be considered. The overall effect caused by temperature changes is hence both variable in magnitude and sign for measurement channels of two similar products as well as for a measurement and reference channel within the same instrument with separate filters, detectors and electronics. For this reason and in order to enhance the measurement quality, the entire NDIR detector is temperature stabilized within the sensor. An active temperature stabilization furthermore helps to reduce the required warm-up time present in any NDIR detector due to self heating effects. A separation of the light source and the detector from the gas stream by windows enhances the temperature stabilization capability and protects the sensitive detector with the filter. A gas heater further reduces the temperature gradient within the cuvette and simultaneously reduces the risk of condensation within the optical components of the sensor.

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Any sensor based on an absorption principle such as an NDIR detector senses the highest raw signal in the case of a complete absence of the target molecules in the beam path as in that case no absorption occurs and the maximum radiation intensity reaches the detector (see (1)). Therefore the regular determination of the sensor signal of a zero gas, the *zeroing*, is essential to account for drift effects that alter the light intensity with time and that otherwise would be erroneously interpreted as changes in target gas concentration. Typical effects are:

(i) Intensity variations or spectral shifts of the light source over time,

305	(ii)	Contamination of any component within the beam path that might cause
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- (iii) Aging effects that alter the detector sensitivity over time,
- (iv) Changes in the pre-amplifier gain of the detector.

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Within the sensor the zeroing does not only account for long-term drift influences but also for changes of the measurement conditions such as large changes in water temperature that cause internal temperature gradients and different water vapor concentrations within the gas stream.

Drift compensation by means of a differential setup in our case realized in the form of filter correlation is referred to as a two-beam/two-wavelength method. It is supposed to compensate any unwanted influences that cause signal drift of both channels in the same manner (aspects (i) and (ii) above) as the measured signal is continuously referenced. Any effects that cause changes in the detector signals and that are not caused by actual concentration changes within the cuvette shall be compensated in real time and parallel to measurements. In reality, this technique has its limitations in accounting for influences resulting from the usage of the two different channels with their own filters and detectors. Theoretically the zero-point of a two-beam instrument should not be affected by the above mentioned drift reasons ((i) through (iv)). But as we measure at two different wavelengths, spectral differences as well as effects related to the two physically different detectors still affect the zero signal of the "two-beam corrected" signal. Therefore we combine the zeroing and the two-beam drift correction means within our sensor. The latter provides a continuous correction applied parallel to measurements, while the zeroing discontinuously further enhances the drift correction capabilities by

correcting for effects that differently affect the measured intensities at both detectors.

Related to the origin of NDIR detector drift in combination with the dual wavelength setup it should be pointed out that changes in the zero concentration measurements can

(i) occur erratically especially after transportation of storage, (ii) are not necessarily linear in time with (iii) the slope commonly decreasing over running time or (iv) even changing its sign.

The zero correction of a two-beam sensor signal does not cover concentration dependent effects that equate to changes in the characteristics of the NDIR sensor's calibration polynomial. Hence for achieving best accuracies particularly with two-beam NDIR sensors, the sensor in addition to regular zeroings needs to be re-calibrated after deployment at different concentrations.

d. Membrane equilibration

The solution-diffusion model can be used to describe gas transport through a dense, semipermeable membrane. Assuming a partial pressure gradient between the water phase and
the internal gas stream, the first step within the transport process is adsorption of the
molecule at the membrane surface. In the case of a sufficiently small gas concentration
the dissolution process into the membrane can be explained by the Henry-Dalton Law,
which states the linearity between the partial pressure of a gas and the concentration of
that gas in solution connected by a temperature, pressure and, in the case of sea water,
salinity dependent constant. The dissolution process is followed by diffusion based
transport of the molecules within the membrane along the concentration gradient.
Outgasing into the headspace again follows the Henry-Dalton Law. This process
continues within the membrane of the sensor until partial pressure equilibrium with

respect to every single gas component is achieved. Hence semi-permeable membranes can be used as equilibrators for dissolved gas measurements (McNeil et al. 2006). They enable the application of gas phase based measuring techniques in an aquatic medium. The time limiting and overall response time dominating step within the entire transport process is the diffusion within the membrane material as well as within a boundary layer in front of the membrane. The permeability of a material is a parameter integrating both, the solubility of a gas within the membrane substance as well as its diffusion constant within the membrane according to Fick's Law. Like the solubility and the diffusion coefficient, the permeability is also temperature and pressure dependent as well as concentration independent. The permeabilities of different gases for a given membrane material differ (Robb 1968; Merkel et al. 2000) and hence determine the time constant for the corresponding partial pressure equilibration. The direction of the transport process is defined by the orientation of the individual partial pressure gradient. Silicone, polydimethylsiloxane (PDMS), was chosen as the membrane material due to its high permeability for CO₂ (Merkel et al. 2000). In addition to all the dissolved gases, water vapor also permeates through the membrane. Its amount within the gas stream is related to the temperature and salinity driven water vapor pressure. The risk of condensation within the headspace and

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membrane. Its amount within the gas stream is related to the temperature and salinity driven water vapor pressure. The risk of condensation within the headspace and especially within the NDIR absorption path induced by steep gradients of warm to cold water is minimized by the heating of the gas on its way to the NDIR detector.

Furthermore the permeability of water vapor is more than a magnitude higher than for CO₂ (Robb 1968), leading to time constants for water vapor equilibration of about 10 s at the given CO₂ response time of about 70 s for this sensor.

Typically silicone layer thicknesses of around 10 µm are used. The thickness is determined during membrane production by permeability measurements. Pure silicone monolayer membranes of this thickness could not be easily handled. Therefore we use thin film composite membranes consisting of the dense silicone layer on top of supporting substructures. In the case of no or minor fouling the membranes can be deployed for several months to years. Cleaning of the membranes with e.g. diluted sulfuric or oxalic acid at pH 2 has successfully been tested. To avoid physical damage of the thin silicone layer mechanical cleaning of the membrane surface should be avoided. Instead the membrane should and can be changed even in the field. The response time of the sensor determined from the flush interval data can also be used to identify organic ongrowth, since heavy fouling slows down the membrane permeation process or the volume rate of pumped water, which both leads to a reduced response time of the sensor.

3. In-water calibration setup

Along with the sensor development we established a laboratory calibration setup for direct underwater pCO_2 calibration (Fig. 3) that can hold up to three sensors simultaneously. An early version of the setup was successfully used in Friedrichs et al. (2010). The setup includes a 120 L insulated and temperature stabilized water tank. The water temperature can be controlled over the temperature range of 0°C to 30°C to within ± 0.02 °C. It is filled with de-ionized water. Sodium carbonate and bicarbonate are added in the required quantities in order to mimic the CO_2 buffer system of seawater (DIC:TA ratio) and thereby allow for a better pCO_2 level control. Silver nitrate is added as an antifoulant. Water is continuously pumped through a reference flow-through pCO_2 system (Körtzinger et al. 1996), which was slightly modified to suit the laboratory conditions.

Additionally it was equipped with drying components to facilitate continuous reference measurements in dry gas. Special care was taken that the return flow of the water from the combined bubble-type/laminar-flow equilibrator into the main tank occurs without flow restrictions and thus without altering the pressure conditions within the equilibrator. Temperature probes within the equilibrator and the main tank are regularly calibrated against a reference probe with an accuracy of ±0.02 °C. The flow-through system is equipped with a bench top IR analyzer (LI-6262 or LI-7000, LI-COR Inc, Lincoln, Nebraska, USA), which is calibrated against 3 primary (certified National Oceanic and Atmospheric Administration (NOAA) standards) or secondary (referenced to NOAA standards) CO₂-in-natural-air standards in the beginning and in the end of each calibration run. Processing of the flow-through system data is carried out according to the procedures described in Dickson et al. (2007) and Pierrot et al. (2009), leading to pCO₂ reference values referred to the water temperature in the tank with an accuracy of 2-3 μatm. During a full calibration run, the *p*CO₂ of the tank water is altered by pH variation through injection of NaOH or HCl solutions. A new concentration in the tank water is set and resolved by the reference system with a time constant of approx. 150 s. By application of the pH-varying technique a wide pCO_2 -range can be realized and calibration steps can be set as desired. Since the overall setup cannot be entirely encapsulated from the surrounding air a small drift of the partial pressures in the tank of typically around 3 µatm hr⁻¹ can be observed. Magnitude and timescale of this drift as well as the fact that both, the flow-through system and the underwater sensor, detect this pCO₂ change make this effect negligible with respect to the assumed accuracy of the calibration process. The adjustable range is limited by the measurement range of the IR-

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analyzer of the flow-through system, which is $3000 \,\mu\text{mol mol}^{-1}$ for the dry CO_2 mole fraction in the equilibrated gas stream (xCO_2). It was found that four calibration steps are sufficient for a sensor calibration in the range of 200 to 1000 μ atm. The course of a calibration is depicted in Fig. 2.

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There are several reasons to calibrate the sensor in water against a proven underway system as opposed to a mere dry gas calibration of the IR detector. The fact that some of the following influences are already considered in data processing corrections or their minimization was addressed in sensor design, does not debilitate the following compensatory advantages of an in-water calibration. Firstly, the temperature stabilization including the gas heater used in the sensor does not completely avert the presence of temperature gradients within the instrument's housing, components and gas stream. Hence the temperature stabilization can only minimize the above mentioned possible temperature effects on NDIR sensors. An in-water calibration at a temperature as close as possible to the expected temperature in the field helps to further reduce these signal influences. Secondly, an in-water calibration as described compensates for all effects related to the large absolute humidity present in the sensor's gas stream. These effects are gas-gas-interactions causing band broadening, potential cross sensitivities of the NDIR signal against H₂O due to minor H₂O absorption at the transmitted wavelengths, or H₂O molecule interaction with the cuvette's surface. For a sensor calibration at only one temperature it is deliberately accepted that the humidity related compensations are only entirely compensated for a deployment at a water temperature equal to the calibration temperature. Thirdly, an in-water calibration compares the fully processed signal of the instrument with a reference value. Hence it characterizes the

overall instrument including the entire membrane equilibration process of the headspace as identified to be important by Byrne et al. (2010). Any not otherwise considered effects occurring in the sensor's gas stream and behind the membrane are taken care of by an inwater calibration.

4. Data processing

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The dual-beam NDIR detector provides two signals. The raw signal, S_{raw} , corresponds to the transmitted light intensity around the wavelength at which CO₂ efficiently absorbs (4.26 μ m) and the reference signal, S_{ref} , expresses the intensity at around 4 μ m where practically no relevant absorption occurs. Water vapor is weak absorber at both wavelengths. A continuously referenced sensor signal, the "two-beam signal", is:

$$S_{\text{2beam}} = \frac{S_{\text{raw}}}{S_{\text{ref}}} \,. \tag{3}$$

As a result of the regular zeroing (Z), we obtain "two-beam zero signals" at discrete points in time:

$$S_{\text{2beam},Z} = \frac{S_{\text{raw},Z}}{S_{\text{ref}\,Z}} \,. \tag{4}$$

- Linear interpolation in time between two adjacent $S_{2\text{beam},Z}$ provides zero reference signals for every point in time: $S_{2\text{beam},Z}(t)$.
- An NDIR signal that is improved by both drift correction (DC) means, dual beam and zeroing, is derived as follows:

$$S_{\rm DC}(t) = \frac{S_{\rm 2beam}(t)}{S_{\rm 2beam,Z}(t)} \tag{5}$$

- The final, drift corrected NDIR signal, which is assumed to be directly
- 460 proportional to the amount of target molecules in the beam path is related to $S_{DC}(t)$.

During calibration the flow-through system provides a reference pCO₂ for every calibration step. The pCO_2 equilibrium is assumed to be established in the membrane equilibrator of the sensor at that time. Using data of the peripheral sensors in the gas stream, the xCO_2 in moist air present at the NDIR detector is derived. As the NDIR signal is proportional to the number of molecules in the beam path and not to xCO_2 , the reference xCO₂ needs to be density corrected by using data of the additional temperature and pressure sensors built into the gas stream of the pCO_2 sensor. A polynomial of rank 3 with a forced zero crossing is then used to calibrate the individual sensor characteristics. It correlates the absorbance signal of the NDIR detector with the corresponding and density corrected xCO_2 in the gas stream (c.f. Fig. 4). Now, all required dependencies are known and the sensor provides the pCO_2 based on the absorbance signal of its NDIR detector in combination with the data of the peripheral sensors in its gas stream. Beside the density correction no other NDIR signal correction addressing a band broadening effect as a consequence of CO₂-H₂O molecule interactions or any other H₂O cross sensitivity is explicitly included in the sensor sided data processing at this point. These aspects are considered through the in-water calibration as mentioned in section 3. Since the consideration of the zeroings requires an interpolation in time, this calculation step is best applied during post processing of field data to obtain a smooth behavior. In order to achieve the best measurement accuracy, changes in the sensor characteristics should also be included into the processing. Therefore an interpolation

between the polynomial of a pre-deployment calibration and the polynomial of a post-

deployment calibration over the course of the deployment is conducted. We apply an

interpolation that is not linear with time, but instead linear with the actual value of the

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zero signal throughout a deployment, $S_{2\text{beam},Z}(t)$. The pre-deployment polynomial is transformed to the post-deployment polynomial by proportionately using the coefficients of the two polynomials according to the actual zero signal. This approach assumes a causal relationship between the temporally often non-linear change in the zero signal and the change in concentration dependent sensor response.

The entire calibration calculations as well as the post processing are accomplished with custom-designed LabVIEW routines (National Instruments, Austin, Texas).

5. Field evaluation

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Field evaluations of the new pCO₂ sensor were carried out in April/May 2010 during a 6week cruise in the North and South Atlantic (R/V Polarstern, ANT-XXVI/4) and in June/July 2011 during a 4-week cruise in the eastern tropical Atlantic (R/V Maria S. Merian, MSM-18/3; Fig. 5). During these cruises, oligotrophic (i.e., subtropical gyres) as well as mesotrophic regions (e.g., continental shelfs, equatorial upwelling) provided a reasonably wide range in pCO_2 (295 to 430 μ atm; c.f. Fig 6. and Fig 7. top panel) and temperature (7.4 to 30.3°C, Fig. 5). Furthermore, strong pCO_2 and temperature gradients were found near hydrographic fronts. Thus, the conditions were ideal for a thorough evaluation of the sensor, which was tested during both cruises in nearly identical underway setups: Seawater, either supplied by the ship's clean seawater supply systems (ANT-XXVI/4) or by a submersible pump installed in the moon pool near the ship's keel (MSM-18/3), was pumped to the laboratory into a thermally insulated flow-through box (80 L volume) at a flow rate of approx. 12 L min⁻¹. Sea surface temperature (SST) and sea surface salinity (SSS) were determined for both cruises at the seawater intake. A sensor package containing the pCO₂ sensor (two units during ANT-XXVI/4, HC1 and

HC2; one unit on MSM-18/3, HC3) and an oxygen optode (Model 3830 or 4330, Aanderaa Data Instruments AS, Bergen, Norway), which also provided the water temperature in the flow-through box with an accuracy of $\pm 0.05^{\circ}$ C, were placed in the underway box. Data were binned into 1-min intervals. Since the pCO₂ sensor data were initially stored at 5-s intervals the transformation to the 1-min intervals represents an averaging of typically 12 spot values. Zeroings were carried out every 12 hours. The membrane interface of the pCO_2 sensor was supplied with a constant seawater flow by a SBE 5T pump (Sea-Bird Electronics Inc., Bellevue, Washington). A fully automated pCO₂ underway instrument (GO, General Oceanics, Miami, USA; Pierrot et al. 2009) based on a spray head equilibrator and a LI-7000 CO₂ analyzer was operated in parallel. Throughout the expeditions and beside the deployment in the flow-through box the sensors were additionally used for measurements on a CTD rosette system (HC1 and HC3; part of the data shown in Fiedler et al. 2013) as well as on a surface drifter (HC3; unpublished data). The two sensors HC1 and HC2 were calibrated at 19.67°C before and after the deployment. HC3 was pre- and post-calibrated at 27.00°C and 26.00°C, respectively. The averaged sensor data were prepared for processing by filtering for obvious outlier zero values as, for example, caused by improper sensor warm-ups, as well other outliers and data biased through insufficient water supply to the flow-through box. All pCO₂ sensor data recorded during flush and zero intervals were excluded from the comparison with the reference. The data of HC2 obtained between 28th of April and 4th of May were excluded from the comparison as well, as the excessively high water temperature did not allow for temperature stabilization at the calibration control

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temperature anymore. Although the control temperature within HC1 also temporarily exceeded the set point, no data were removed here, as the unit seemed to be more robust in this respect as compared to HC2.

Figures 6 and 7 show the pCO_2 measured by the flow-through reference system as well as the pCO_2 data of the HC1-3 sensors with all datasets corrected to SST for direct comparison (Takahashi et al. 1993). In addition, the pCO_2 differences (ΔpCO_2) between the sensor and the reference ($pCO_{2,GO}$) are plotted vs. time. The three ΔpCO_2 curves for every sensor are related to different drift compensation mechanisms:

- (i) pre-calibration, no zeroings
- (ii) pre-calibration + zeroings

540 (iii) pre-calibration + zeroings + post-calibration

In the case of (i) only the dual-beam correction is applied and the polynomial of the pre-deployment calibration is used for the entire data set. Since in this case data are processed by referencing all signals to the zeroing carried out during shore-based calibration, large offsets can already occur at the beginning of a deployment due to a sensor drift that occurred during storage and transport or due to measurement conditions that strongly deviate from calibration conditions. Also more or less strong drift behavior is observed over time. When the regular zero correction is carried out as described above (section 4), both initial offsets and drift over the course of the deployment are strongly reduced demonstrating the effectiveness of this first correction. However, even smaller residuals were obtained when both zeroings as well as pre- and the post-deployment calibration polynomials are considered in the post-processing (section 4). Table 2 shows the statistics of the pCO_2 residuals as obtained by this optimized processing routine. In all

three sensor deployments the mean $p\text{CO}_2$ offset to the reference system is within $\pm 3~\mu$ atm. With an average $p\text{CO}_2$ residual over all three sensors of -0.6 \pm 3.0 μ atm (RMSE=3.7 μ atm) no systematic offset between $p\text{CO}_2$ sensor and reference system could be found. This indicates that the agreement between sensor and flow-through system $p\text{CO}_2$ is of similar magnitude as the accuracy of the flow-through system which was estimated at 2 μ atm (Pierrot et al. 2009). We use the mean RMSE from all three deployments of 3.7 μ atm as a conservative estimate of the sensor accuracy. This result is very promising as the sensors were only calibrated at a single water temperature and experienced a large temperature range during deployment (>20°C). We note that during these field tests the sensors ran autonomously and without maintenance but were also used in other tests (e.g., deployments of CTD rosette casts; part of the data shown in Fiedler et al., 2013). For this purpose the instruments had to be restarted several times. This appears to not have affected sensor performance negatively, since sensor HC3 was restarted most frequently but shows the smallest overall offset.

The compensation routines applied, which exclusively rely on data measured by the sensors themselves and the information obtained from calibrations, account well for the signal drift for all 3 sensors. After processing the data do not exhibit significant unaccounted drift behavior. This is even the case for the most strongly drifting sensor HC1, whose signal change over time is also reflected by the change in its calibration polynomials (Fig. 4). Although an NDIR detector drifting as strongly as in the case of HC1 would not pass current quality controls within the manufacturer's production, it is still a good example to demonstrate the effectiveness of the described processing algorithms. We would like to note that sensor HC1 also participated in a different sensor

evaluation project (Tamburri et al. 2011). At that time, the sensor also showed a strong drift that could not be adequately compensated for due to lack of the full understanding of the required post-processing steps that is presented here. Although the drift of HC1 could be well corrected for, in the end it still shows a slightly larger mean value and RMSE compared to HC2 and HC3. NDIR detectors that show a smaller zero drift, typically show a smaller concentration dependent signal change over time as well. This conclusion is further corroborated by the observation that a transformation of the pre-deployment into the post-deployment calibration polynomial based on the course of the zero values finally provides a better correction as opposed to a transformation assumed to occur linear in time (data not shown here).

In order to identify any remaining issues in the drift corrected pCO_2 sensor data, the ΔpCO_2 residuals were plotted against $pCO_{2,GO}$, SST and pH_2O (Fig. 8). All sensors show a weak correlation with all three parameters $(0.0 < R^2 < 0.6)$, which is most clear for sensor HC2 $(0.5 < R^2 < 0.6)$. Since $pCO_{2,GO}$, SST and pH_2O are strongly correlated in the field data, the cause of these remaining residual correlations cannot be clearly discerned. Nevertheless, there is indication for a weak NDIR signal dependency on pH_2O . In fact, CO_2 -NDIR detectors over-estimate in the presence of water vapor due to pressure broadening effects (McDermitt et al. 1993). Furthermore, the magnitude of this effect increases with pCO_2 and pH_2O . Since the sensor calibrations were only conducted at one temperature and band broadening effects due to varying water vapor concentrations are not considered within the sensor's data processing, this might explain at least part of the dependencies. The fact that HC2 and HC3 show the smallest ΔpCO_2 around the water temperature at which they were calibrated, support this observation. The limitations of

the present data set do not allow to further investigate this issue and additional tests have to be carried out to assess the potential for further improvement, e.g., by performing laboratory tests with and calibrations of the pCO_2 sensor at more than one temperature.

6. Summary and outlook

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The development of a new underwater pCO_2 sensor based on membrane equilibration and NDIR spectrometry was described. Special emphasis was put on compensation measures for NDIR sensor drift as well as on the in-water calibration of the sensor. The performance of the pCO_2 sensor was assessed based on surface water field data obtained during two cruises both lasting at least one month and covering a wide range in pCO_2 (289 to 445 µatm) and temperature (7.4 to 30.1°C). A wet gas stream within a small underwater sensor represents a demanding environment for NDIR detectors. Against this background the observed mean offset of -0.6 ± 3.0 µatm with a RMSE of 3.7 µatm to the pCO₂ reference instrument as obtained through application of a pre- and post-deployment calibration in combination with regular zeroings is a very promising result, especially since the sensors were also used for various other measurements including profiling applications in the water column during the deployments (data not part of this assessment). At a pCO_2 of 400 μ atm the observed mean pCO_2 difference corresponds to about 0.2 % and the RMSE to less than 1 %. This favorable result underlines the efficiency of the applied processing algorithms. The acquisition of high-quality field data by the new pCO_2 sensor calls for regular checks of the sensor parameters gas temperature, pressure, RH and control temperature to guarantee optimal functioning of the sensor. In addition, regular zero gas measurements need to be carried out at least every 12 h under deployment conditions as a drift correction means beside the inherent

single-beam dual wavelength setup of the sensor. To properly apply the zero information during post processing measurement data should always be flanked by zeroings. Thirdly, in order to also account for the concentration dependant effects on zero and dual beam corrected signals, the sensor needs to be re-calibrated at different pCO_2 levels on a time scale of several months to a year to achieve the highest accuracies through data post processing. A water calibration as presented in this paper at a temperature close to the expected water temperatures in the field is beneficial. Nevertheless, the RMSE found within this work is based on field data obtained under conditions where water temperatures deviated by $\pm 10^{\circ}C$ from the calibration temperature.

The assessment given here represents an important milestone for the development of the sensor. The procedures discussed are planned to be further automated and implemented into data processing routines. The next development steps include investigation of potential improvements with respect to NDIR data processing and the laboratory calibration routines to identify and compensate for minor signal dependencies on water vapor and on changes in gas matrix composition as caused by e.g. strongly varying oxygen concentrations. Moreover the long-term stability during deployments on moorings and profiling buoys in the water column will be investigated as well as the sensor performance on different moving platforms analyzed. Especially the latter has been simplified by a recently released smaller and faster version of the HydroC.

Acknowledgments.

The authors would like to thank all members of the "(C)O₂" working group at the GEOMAR as well as Gernot Friedrichs from the CAU for fruitful discussions and helpful advice. Furthermore, we thank the CONTROS team for their support in this development project. We explicitly acknowledge Christian Rettich for his assistance especially during the construction of the first prototypes as well as Claus Hinz and Matthias Lunge for valuable help with laboratory experiments. Janna Bohlen and Christoph Kirbach are thanked for commenting on the manuscript. We would also like to thank captains and crews of R/V *Polarstern* and of *Maria S. Merian* for assistance. This work was partly funded by the SAW project OCEANET of the Leibniz Association (2008-2010) and by the SOPRAN project of the German Federal Ministry of Education and Research (grants 03F0462A and 03F0611A).

References

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Byrne, R. H., and W. Yao, 2008: Procedures for measurement of carbonate ion 656 657 concentrations in seawater by direct spectrophotometric observations of Pb(II) 658 complexation. *Mar. Chem.*, **112**, 128–135, doi:10.1016/j.marchem.2008.07.009. 659 Byrne, R. H., X. Liu, and E. A. Kaltenbacher, 2002: Spectrophotometric measurement of 660 total inorganic carbon in aqueous solutions using a liquid core waveguide. Anal. Chim. Acta, 451, 221–229, doi:10.1016/S0003-2670(01)01423-4. 661 662 Byrne, R. H. and Coauthors, 2010: Sensors and Systems for in situ Observations of 663 Marine Carbon Dioxide System Variables. Proceedings of OceanObs'09: Sustained 664 Ocean Observations and Information for Society, J. Hall, D.E. Harrison, and D. 665 Stammer, Eds., Venice, Italy, ESA Publication WPP-306. 666 DeGrandpre, M. D., T. R. Hammar, S. P. Smith, and F. L. Sayles, 1995: In-Situ Measurements of Seawater pCO₂. Limnol. Oceanogr., **40**, 969–975. 667 Dickson, A. G., C. L. Sabine, and J. R. Christian, eds., 2007: Guide to best practices for 668 669 ocean CO₂ measurements. PICES Special Publication 3. 670 Doney, S. C., V. J. Fabry, R. A. Feely, and J. A. Kleypas, 2009: Ocean Acidification: The 671 Other CO₂ Problem. Ann. Rev. Mar. Sci., 1, 169–192, 672 doi:10.1146/annurev.marine.010908.163834. 673 Fiedler, B., P. Fietzek, N. Vieira, P. Silva, H. C. Bittig, and A. Körtzinger, 2013: In situ 674 CO₂ and O₂ measurements on a profiling float. J. Atmos. Oceanic Technol., 30, 112– 675 126. Fietzek, P., S. Kramer, and D. Esser, 2011: Deployment of the HydroCTM (CO₂/CH₂) on 676 stationary and mobile platforms - Merging the trends in the field of platform and 677 678 sensor development. Proceedings of Oceans 11 MTS/IEEE Conference, 19-22 Sept. 679 2011, Kona, Hawaii, USA, 1-9. 680 Friedrichs, G., J. Bock, F. Temps, P. Fietzek, A. Körtzinger, and D. W. R. Wallace, 2010: Toward continuous monitoring of seawater $^{13}CO_2/^{12}CO_2$ isotope ratio and pCO_2 : 681 Performance of cavity ringdown spectroscopy and gas matrix effects. *Limnol*. 682 683 Oceanogr.: Methods, 8, 539–551, doi:10.4319/lom.2010.8.539. 684 Gruber, N. and Coauthors, 2010: Toward an Integrated Observing System for Ocean 685 Carbon and Biogeochemistry at a Time of Change. *Proceedings of OceanObs'09*: 686 Sustained Ocean Observations and Information for Society, J. Hall, D.E. Harrison,

and D. Stammer, Eds., Venice, Italy, ESA Publication WPP-306.

- 688 IPCC, 2005: IPCC special report on carbon dioxide capture and storage. B. Metz, O.
- Davidson, H. De Coninck, M. Loos, and L. Meyer, Eds. Cambridge University
- 690 Press.
- Johnson, K. S. and Coauthors, 2009: Observing biogeochemical cycles at global scales
- with profiling floats and gliders: prospects for a global array. Oceanography, 22,
- 693 216–224.
- Körtzinger, A., H. Thomas, B. Schneider, N. Gronau, L. Mintrop, and J. C. Duinker,
- 695 1996: At-sea intercomparison of two newly designed underway pCO₂ systems -
- 696 encouraging results. *Mar. Chem.*, **52**, 133–145.
- Körtzinger, A. and Coauthors, 2000: The international at-sea intercomparison of fCO₂
- systems during the RV Meteor Cruise 36/1 in the North Atlantic Ocean. *Mar.*
- 699 *Chem.*, **72**, 171–192, doi:10.1016/S0304-4203(00)00080-3.
- Lefèvre, N., J. P. Ciabrini, G. Michard, B. Brient, M. DuChaffaut, and L. Merlivat, 1993:
- A new optical sensor for pCO₂ measurements in seawater. Mar. Chem., 42, 189–
- 702 198, doi:10.1016/0304-4203(93)90011-C.
- Martz, T. R., J. G. Connery, and K. S. Johnson, 2010: Testing the Honeywell Durafet for
- seawater pH applications. *Limnology And Oceanography Methods*, **8**, 172–184,
- 705 doi:10.4319/lom.2010.8.172.
- McDermitt, D. K., J. M. Welles, and R. D. Eckles, 1993: Effects of temperature, pressure
- and water vapor on gas phase infrared absorption by CO₂. *LI-COR*, *Inc*, *Lincoln*,
- 708 *USA*., 1–5.
- McNeil, C., E. D'Asaro, B. Johnson, and M. Horn, 2006: A Gas Tension Device with
- Response Times of Minutes. J. Atmos. Oceanic Technol., 23, 1539,
- 711 doi:10.1175/JTECH1974.1.
- Merkel, T. C., V. I. Bondar, K. Nagai, B. D. Freeman, and I. Pinnau, 2000: Gas sorption,
- diffusion, and permeation in poly(dimethylsiloxane). J. Polym. Sci., Part B: Polym.
- 714 *Phys.*, **38**, 415–434, doi:10.1002/(SICI)1099-0488(20000201)38:3<415::AID-
- 715 POLB8>3.0.CO;2-Z.
- Millero, F. J., 2007: The marine inorganic carbon cycle. *Chem. Rev.*, **107**, 308–341,
- 717 doi:10.1021/cr0503557.
- Pfeil, B. and Coauthors, 2013: A uniform, quality controlled Surface Ocean CO₂ Atlas
- 719 (SOCAT). Earth Syst. Sci. Data, 125–143, doi:10.5194/essd-5-125-2013.
- Pierrot, D. and Coauthors, 2009: Recommendations for autonomous underway pCO₂
- measuring systems and data-reduction routines. *Deep-Sea Res.*, *Part II*, **56**, 512–
- 722 522, doi:10.1016/j.dsr2.2008.12.005.

- Robb, W. L., 1968: Thin silicone membranes Their permeation properties ans some applications. *Annals New York Academy of Science*, 119–137.
- Roemmich, D., G. C. Johnson, S. Riser, R. Davis, and J. Gilson, 2009: The Argo
- Program: observing the global ocean with profiling floats. *Oceanography*, **22**, 34–
- 727 43.
- Rogner, H. H., D. Zhou, R. Bradley, P. Crabbé, O. Edenhofer, B. Hare (Australia), L.
- Kuijpers, and M. Yamaguchi, 2007: *Introduction. In Climate Change* 2007:
- 730 Mitigation. Contribution of Working Group III to the Fourth Assessment Report of
- 731 *the Intergovernmental Panel on Climate Change*. Cambridge University Press.
- Sabine, C. L. and Coauthors, 2004: The Oceanic Sink for Anthropogenic CO₂. Science,
- 733 **305**, 367–371.
- Saderne, V., P. Fietzek, and P. M. J. Herman, 2013: Extreme variations of pCO₂ and pH
- in a macrophyte meadow of the Baltic Sea in summer: evidence of the effect of
- photosynthesis and local upwelling. *PLoS ONE*, **8**, e62689,
- 737 doi:10.1371/journal.pone.0062689.
- Seidel, M. P., M. D. DeGrandpre, and A. G. Dickson, 2008: A sensor for in situ
- indicator-based measurements of seawater pH. Mar. Chem., 109, 18–28,
- 740 doi:10.1016/j.marchem.2007.11.013.
- Severinghaus, J. W., and A. F. Bradley, 1958: Electrodes for Blood pO₂ and pCO₂
- 742 Determination. *J. Appl. Physiol.*, **13**, 515–520.
- Stow, R. W., R. F. Baer, and B. F. Randall, 1957: Rapid measurement of the tension of
- carbon dioxide in blood. Archives of physical medicine and rehabilitation, **38**, 646–
- 745 650.
- Takahashi, T., 1961: Carbon dioxide in the atmosphere and in Atlantic Ocean water. *J.*
- 747 *Geophys. Res.*, **66**, 477–494.
- Takahashi, T., J. Olafsson, J. G. Goddard, D. W. Chipman, and S. C. Sutherland, 1993:
- Seasonal variation of CO₂ and nutrients in the high-latitude surface oceans: A
- 750 comparative study. *Global Biogeochem. Cycles*, **7**, 843–878.
- 751 Takahashi, T. and Coauthors, 2009: Climatological mean and decadal change in surface
- ocean pCO₂, and net sea-air CO₂ flux over the global oceans. Deep-Sea Res., Part
- 753 *II*, **56**, 554–577, doi:10.1016/j.dsr2.2008.12.009.
- 754 Tamburri, M. N. and Coauthors, 2011: Alliance for Coastal Technologies: Advancing
- Moored pCO₂ Instruments in Coastal Waters. Mar. Technol. Soc. J., 45, 43–51,
- 756 doi:10.4031/MTSJ.45.1.4.

757 758	Wang, Z. A., X. Liu, R. H. Byrne, R. Wanninkhof, R. E. Bernstein, E. A. Kaltenbacher, and J. Patten, 2007: Simultaneous spectrophotometric flow-through measurements
759 760	of pH, carbon dioxide fugacity, and total inorganic carbon in seawater. <i>Anal. Chim. Acta</i> , 596 , 23–36.
761 762	Watson, A. J. and Coauthors, 2009: Tracking the variable North Atlantic sink for atmospheric CO ₂ . <i>Science</i> , 326 , 1391–1393, doi:10.1126/science.1177394.
763 764	Weiss, R. F., 1974: Carbon dioxide in water and seawater: the solubility of a non-ideal gas. <i>Mar. Chem.</i> , 2 , 203–215.
765 766 767	Wiegleb, G. and Coauthors, 2001: <i>Industrielle Gasssensorik: Messverfahren - Signalverarbeitung - Anwendungstechnik - Prüfkriterien</i> . G. Wiegleb, Ed. expert-Verlag, Renningen.
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List of Figures

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771 TABLE 1. Specifications of the developed pCO_2 sensor as used during the deployments 772 discussed in this paper. The specifications of the currently available sensor model differ 773 from these values with respect to size and power consumption. The power required for 774 the temperature stabilization as well as the warm-up duration depend on the actual water 775 temperature, the chosen control temperature as well as on the thickness of the insulation 776 material. The given warm-up times correspond to 24 V supply voltage in 20°C water and 777 to 12 V supply voltage in 3°C water. Please refer to the text for further details regarding 778 the warm-up and the zeroing interval. The response times refer to the usage of two 779 different pump models at 20°C water temperature. The pump SBE 5T has a flow rate of approx. 105 mL s⁻¹, while the smaller model, SBE 5M, provides a flow rate of approx. 780 35 mL s^{-1} . 781 782 783 TABLE 2. Statistics of the $\Delta p \text{CO}_2$ residuals for all three field deployments of the $p \text{CO}_2$ 784 sensor with mean, standard deviation, σ , and root mean square error, RMSE, for a total of 785 *n* observations. Also shown is the mean of all three deployments. 786 787 FIG. 1. Drawing of the pCO_2 sensor as used in the present study (upper panel). The sensor 788 is equipped with a water pump (SBE 5T) and a flow-head. The lower panel shows a 789 schematic drawing of the sensor. Partial pressure equilibration occurs at the planar, semi-790 permeable membrane separating the water from the internal head space of approx. 20 mL. 791 A pump continuously circulates the gas between the membrane equilibrator, a heater and 792 the NDIR detector. Valves can be toggled to realize a zero gas measurement by guiding

the gas stream through a soda lime cartridge instead of through the membrane equilibrator.

FIG. 2. Course of a calibration carried out at a constant water temperature within the calibration setup described in the text. The absorbance NDIR signal shown in arbitrary units is dual-beam and zero corrected. Also shown is the CO_2 mole fraction (xCO_2) measured by the reference flow-through system. The numbers 1-4 indicate the calibrations steps and "Z", "F" and "M" mark the sensor intervals: zero, flush and measure.

FIG. 3. In-water pCO_2 sensor calibration setup. One water pump is used to allow for sufficient mixing of the water in the tank, dispersion of the injected acid and base as well as to provide water to the reference flow-through system, which is installed in a bypass and whose NDIR unit is regularly calibrated with reference gases.

Fig. 4. Calibration polynomials of sensor HC1 before and after the deployment on R/V *Polarstern*. The absorbance NDIR signal is calibrated against the xCO_2 within the sensor's gas stream obtained according to the text. The polynomials match well the sensor's response characteristics (R^2 in both cases >0.999). A concentration dependent change in the sensor characteristics between the different points in time (i.e., before and after deployment) of the calibrations is clearly visible in this example for a strongly drifting NDIR sensor.

816 FIG. 5. The cruise track of R/V *Polarstern* cruise ANT-XXVI/4 and of R/V *Maria S*. 817 Merian cruise MSM-18/3 (left) as well as the encountered sea surface salinity and 818 temperature (right panel). R/V *Polarstern* sailed from Germany to Chile and R/V *Maria* 819 S. Merian from Cape Verde to Gabon. 820 821 Fig. 6. pCO₂ data obtained by the GO-system as well as two HydroC sensors, HC1 and 822 HC2, during ANT-XXVI/4. The top panel shows the absolute values, while the middle 823 and bottom panel show the differences between the HydroC-pCO₂ and the reference 824 $(\Delta p CO_2)$ over time for three different processing methods in order to visualize the 825 potential of different drift compensation mechanisms (refer to text). The black curve of 826 HC1 and HC2 represent the final drift corrected and post-processed signal. Refer to text 827 for further details on the different processing methods as well as on the peculiarities of 828 strongly drifting HC1. 829 830 Fig. 7. The pCO_2 trace during MSM-18/3 measured by the GO reference pCO_2 system, 831 and a HydroC sensor, HC3 (top panel). The lower panel shows the pCO_2 residuals 832 between the HydroC and the reference (Δp CO₂) for three different drift compensation 833 scenarios. The black curve in the lower panel represents the final drift corrected and post-834 processed signal. 835 836 FIG. 8. In the top row the residuals of HC1 are plotted over the reference pCO₂ (left), the 837 SST (middle) and over the pH₂O as measured within the internal gas stream close to the 838 NDIR detector (right). The middle and the bottom row show the same parameter plots for

839	HC2 and HC3, respectively. Although not statistically significant in all cases, a general
840	positive correlation with all three parameters can be made out and is further discussed in
841	the text.

TABLE 1. Specifications of the developed *p*CO₂ sensor as used during the deployments discussed in this paper. The specifications of the currently available sensor model differ from these values with respect to size and power consumption. The power required for the temperature stabilization as well as the warm-up duration depend on the actual water temperature, the chosen control temperature as well as on the thickness of the insulation material. The given warm-up times correspond to 24 V supply voltage in 20°C water and to 12 V supply voltage in 3°C water. Please refer to the text for further details regarding the warm-up and the zeroing interval. The response times refer to the usage of two different pump models at 20°C water temperature. The pump SBE 5T has a flow rate of approx. 105 mL s⁻¹, while the smaller model, SBE 5M, provides a flow rate of approx. 35 mL s⁻¹.

Measurement principle	IR absorption measurement in a membrane equilibrated			
	headspace			
Detector	Single beam dual wavelength NDIR detector; zeroings at			
	desired intervals			
Housing, dimensions	Cylindrical titanium housing, 90 x 530 mm (without connector)			
Depth capability	2000 m (standard)			
Weight	5.5 kg in air, approx. 2.6 kg in water			
Operating temperature	3 – 30°C			
Supply voltage	11-24 VDC			
Power requirements	<3 W for the detector and all the electronics,			
	+<1 W -3.5 W for temperature stabilization,			
	+ <6 W during warm-up for 2 – 30 min,			

	+ 4 W during zeroing, + water pump: 7 W or 1.5 W (pump SBE 5T and 5M,
	respectively)
Sampling rate	≤ 1 Hz
Response time ($t_{63\%}$)	Approx. 70 s or 130 s (pump SBE 5T and 5M, respectively)
Measurement range	200 μatm – 1000 μatm (standard)

TABLE 2. Statistics of the $\Delta p \text{CO}_2$ residuals for all three field deployments of the $p \text{CO}_2$ sensor with mean, standard deviation, σ , and root mean square error, RMSE, for a total of n observations. Also shown is the mean of all three deployments.

Sensor	Mean (µatm)	σ (µatm)	RMSE (µatm)	n
TIC1	2.1	2.0	4.2	0.4701
HC1	-3.1	2.9	4.2	24791
HC2	1.8	3.4	3.9	24163
HC3	-0.7	2.8	2.8	12770
Overall mean	-0.6	3.0	3.7	

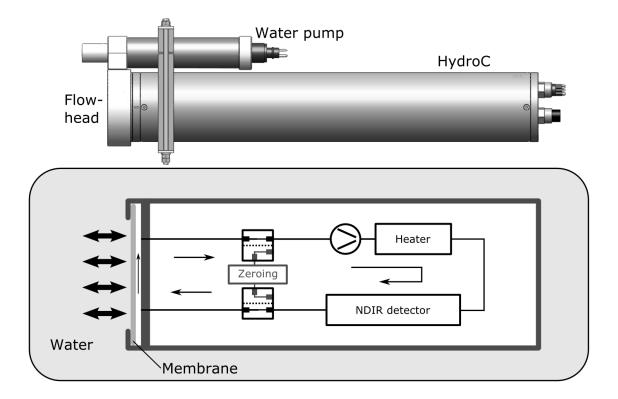


FIG. 1. Drawing of the pCO_2 sensor as used in the present study (upper panel). The sensor is equipped with a water pump (SBE 5T) and a flow-head. The lower panel shows a schematic drawing of the sensor. Partial pressure equilibration occurs at the planar, semi-permeable membrane separating the water from the internal head space of approx. 20 mL. A pump continuously circulates the gas between the membrane equilibrator, a heater and the NDIR detector. Valves can be toggled to realize a zero gas measurement by guiding the gas stream through a soda lime cartridge instead of through the membrane equilibrator.

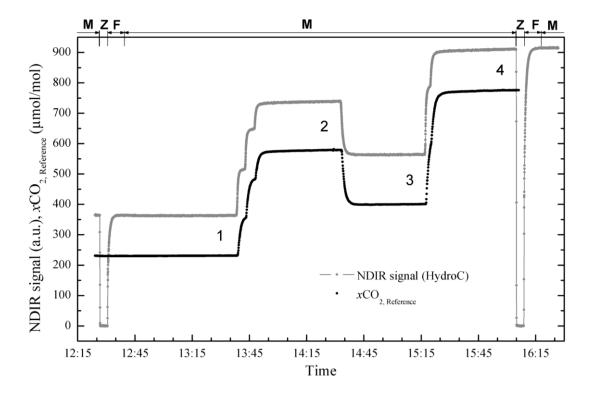


Fig. 2. Course of a calibration carried out at a constant water temperature within the calibration setup described in the text. The absorbance NDIR signal shown in arbitrary units is dual-beam and zero corrected. Also shown is the CO_2 mole fraction (xCO_2) measured by the reference flow-through system. The numbers 1-4 indicate the calibrations steps and "Z", "F" and "M" mark the sensor intervals: zero, flush and measure.

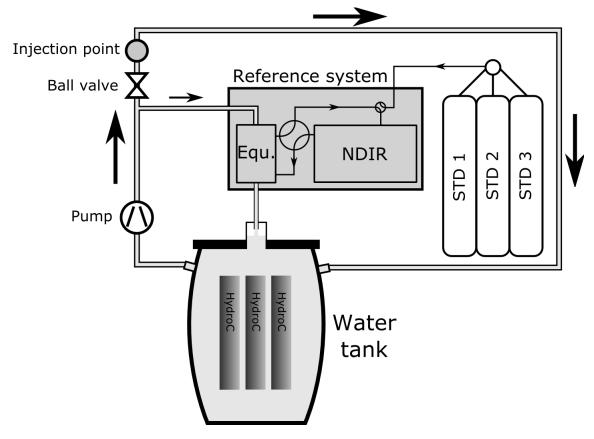


FIG. 3. In-water pCO_2 sensor calibration setup. One water pump is used to allow for sufficient mixing of the water in the tank, dispersion of the injected acid and base as well as to provide water to the reference flow-through system, which is installed in a bypass and whose NDIR unit is regularly calibrated with reference gases.

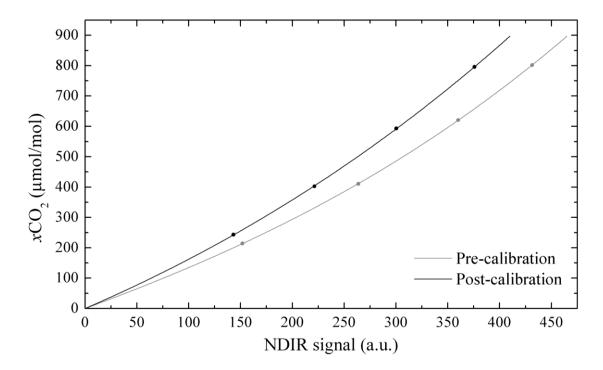


FIG. 4. Calibration polynomials of sensor HC1 before and after the deployment on R/V *Polarstern*. The absorbance NDIR signal is calibrated against the xCO $_2$ within the sensor's gas stream obtained according to the text. The polynomials match well the sensor's response characteristics (R^2 in both cases >0.999). A concentration dependent change in the sensor characteristics between the different points in time (i.e., before and after deployment) of the calibrations is clearly visible in this example for a strongly drifting NDIR sensor.

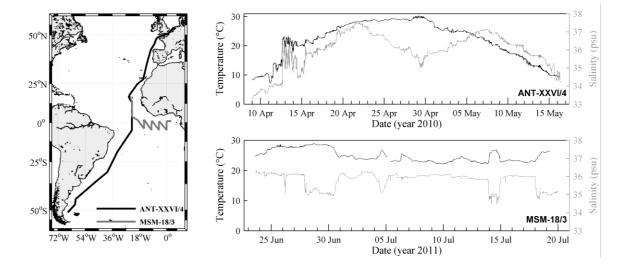


FIG. 5. The cruise track of R/V *Polarstern* cruise ANT-XXVI/4 and of R/V *Maria S*.

Merian cruise MSM-18/3 (left) as well as the encountered sea surface salinity and temperature (right panel). R/V *Polarstern* sailed from Germany to Chile and R/V Maria S. Merian from Cape Verde to Gabon.

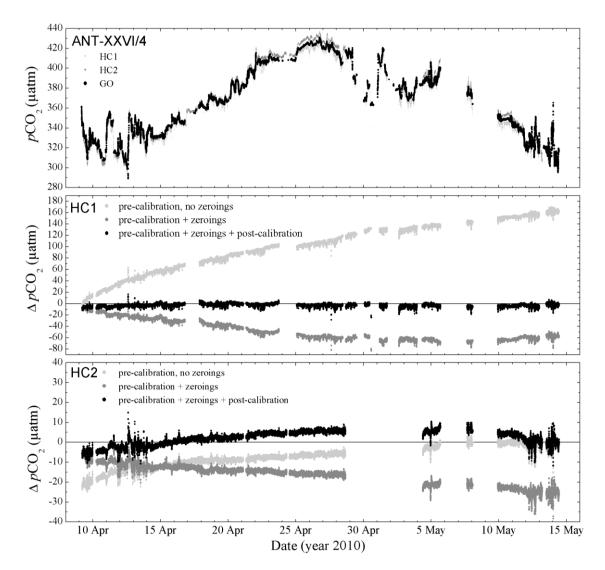


Fig. 6. pCO_2 data obtained by the GO-system as well as two HydroC sensors, HC1 and HC2, during ANT-XXVI/4. The top panel shows the absolute values, while the middle and bottom panel show the differences between the HydroC- pCO_2 and the reference (ΔpCO_2) over time for three different processing methods in order to visualize the potential of different drift compensation mechanisms. The black curve of HC1 and HC2 represent the final drift corrected and post-processed signal. Refer to text for further details on the different processing methods as well as on the peculiarities of strongly drifting HC1.

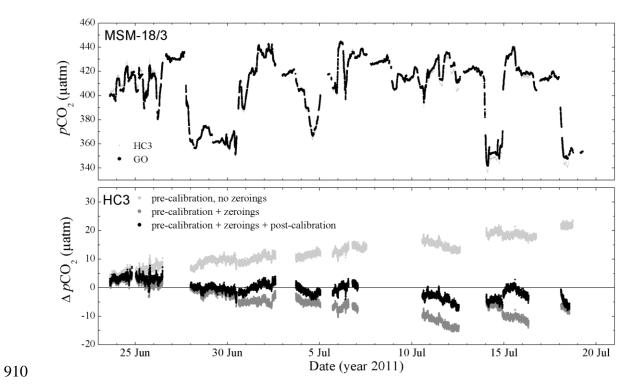


FIG. 7. The pCO_2 trace during MSM-18/3 measured by the GO reference pCO_2 system, and a HydroC sensor, HC3 (top panel). The lower panel shows the pCO_2 residuals between the HydroC and the reference (ΔpCO_2) for three different drift compensation scenarios. The black curve in the lower panel represents the final drift corrected and post-processed signal.

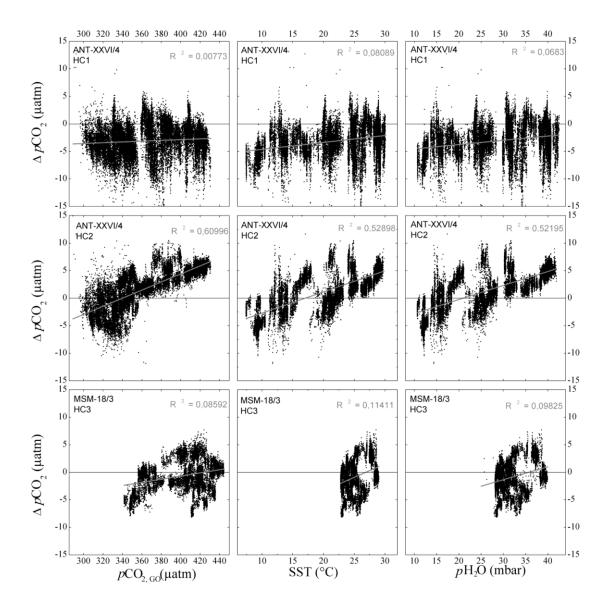


FIG. 8. In the top row the residuals of HC1 are plotted over the reference pCO_2 (left), the SST (middle) and over the pH_2O as measured within the internal gas stream close to the NDIR detector (right). The middle and the bottom row show the same parameter plots for HC2 and HC3, respectively. Although not statistically significant in all cases, a general positive correlation with all three parameters can be made out and is further discussed in the text.