

## Accepted Manuscript

Predictions of critical transitions with non-stationary reduced order models

Christian L.E. Franzke

PII: S0167-2789(13)00221-2

DOI: <http://dx.doi.org/10.1016/j.physd.2013.07.013>

Reference: PHYSD 31414

To appear in: *Physica D*

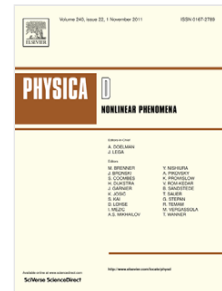
Received date: 15 January 2013

Revised date: 13 July 2013

Accepted date: 18 July 2013

Please cite this article as: C.L.E. Franzke, Predictions of critical transitions with non-stationary reduced order models, *Physica D* (2013), <http://dx.doi.org/10.1016/j.physd.2013.07.013>

This is a PDF file of an unedited manuscript that has been accepted for publication. As a service to our customers we are providing this early version of the manuscript. The manuscript will undergo copyediting, typesetting, and review of the resulting proof before it is published in its final form. Please note that during the production process errors may be discovered which could affect the content, and all legal disclaimers that apply to the journal pertain.



Test of tipping point predictions with reduced order stochastic models

Ensemble approach for tipping point prediction

Ensemble approach is robust in predicting tipping points

ACCEPTED MANUSCRIPT

# Predictions of Critical Transitions with Non-Stationary Reduced Order Models

Christian L. E. Franzke

*British Antarctic Survey, Natural Environment Research Council, Cambridge, UK*

*Phone: +44 (0) 1223221350*

*Fax: +44 (0) 1223221226*

*E-mail: christian.franzke@gmail.com*

---

## Abstract

Here we demonstrate the ability of stochastic reduced order models to predict the statistics of non-stationary systems undergoing critical transitions. First, we show that the reduced order models are able to accurately predict the autocorrelation function and probability density functions (PDF) of higher dimensional systems with time-dependent slow forcing of either the resolved or unresolved modes. Second, we demonstrate that whether the system tips early or repeatedly jumps between the two equilibrium points (flickering) depends on the strength of the coupling between the resolved and unresolved modes and the time scale separation. Both kinds of behaviour have been found to precede critical transitions in earlier studies. Furthermore, we demonstrate that the reduced order models are also able to predict the timing of critical transitions. The skill of various proposed tipping indicators are discussed.

*Keywords:* Stochastic Modeling, Tipping Points, Model Reduction, Non-Stationarity, Bifurcation, Critical Transition

---

## 1. Introduction

Many complex dynamical systems exhibit so-called critical transition or tipping points in which the system approaches a bifurcation point which can lead to sudden and possibly irreversible changes. Even small changes in the control parameter or forcing can lead to a large jump to a different state with

6 possibly catastrophic outcomes. Examples of such tipping points in the real  
7 world range from epileptic seizures (McSharry et al., 2003), financial market  
8 failures (Sornette and Johansen, 1997), ecosystems (Scheffer et al., 2001),  
9 fisheries (Biggs et al., 2009), abrupt shifts in ocean circulation (Monahan et  
10 al., 2008), paleoclimatic abrupt changes (Dakos et al., 2008; Lenton et al.,  
11 2008; Livina et al., 2011, 2013; Cimatoribus et al., 2013), irreversible decline  
12 of the Greenland Ice Sheet (Ridley et al., 2010) and loss of Arctic sea ice  
13 extent (Eisenmann and Wettlaufer, 2008; Wadhams, 2012). As these tipping  
14 points directly affect human well being and the economy it is of utmost  
15 importance to be able to forecast these sudden shifts in order to either avert  
16 them or at least mitigate their effects. Thus, the detection of early warning  
17 signals of imminent tipping points has attracted a lot of attention (Scheffer et  
18 al., 2009; Biggs et al., 2009; Dakos et al., 2008; Ditlevsen and Johnsen, 2010;  
19 Held and Kleinen, 2004; Kuehn, 2011, 2013; Lenton et al., 2008; Sieber and  
20 Thompson, 2012; Livina and Lenton, 2007; Livina et al., 2010, 2011, 2012).

21 Most tipping point detection methods are based on the theory of critical  
22 transitions and critical slowing down. Typical signs of an imminent tipping  
23 point are that the intrinsic transient response to perturbations slows down  
24 (Wissel, 1984; Held and Kleinen, 2004; Veraart et al., 2012), an increase in  
25 autocorrelation (Scheffer et al., 2009), an increase in variance (Ditlevsen and  
26 Johnsen, 2010) or in skewness (Guttal and Jayaprakash, 2008).

27 The slowing down is usually detected by computing the lag-1 autocorre-  
28 lation value by using a sliding data window (Held and Kleinen, 2004) or by  
29 a so-called DFA propagator (Livina and Lenton, 2007). Evidence for an im-  
30 minent tipping point is found when one of the indicators shows an increasing  
31 trend. Unfortunately this approach is sensitive to the used window length  
32 and the detrending procedure before the indicators are computed. Further-  
33 more, the underlying assumption of the detrending procedure is that the time  
34 series over the window length can be considered to be stationary, which is  
35 contradictory to the original assumption that the system is approaching a  
36 bifurcation (Boettinger and Hastings, 2012b). Furthermore, there is also no  
37 real threshold value which needs to be crossed in order to signal that the  
38 system approaches the tipping point. Boettinger and Hastings (2012a) argue  
39 that most previous studies might be biased because they focus only on pe-  
40 riods with a critical transition. They suggest that model based approaches,  
41 especially ensemble predictions, are less subject to this bias.

42 The above proposed tipping point indicators are all based on the analysis  
43 of observed time series. Since in most cases tipping points are singular events

44 the ensemble approach is not feasible with just observational data; though  
45 Cimatoribus et al. (2013) used the Dansgaard-Oeschger events encoded in  
46 Greenland ice cores in an ensemble sense. However, in many situations tip-  
47 ping points are singular events which might not have happened or been ob-  
48 served before. An alternative approach is to use low-order dynamical models  
49 fitted to the observed data to predict tipping points (Carpenter and Brock,  
50 2011). Here we will evaluate the possible use of reduced order stochastic  
51 models in predicting tipping points using an ensemble approach.

52 In a series of papers Majda et al. (1999, 2001, 2002, 2005, 2008, 2009)  
53 developed a systematic framework for the derivation of physics constrained  
54 reduced order models which are nonlinear and have state-dependent noise.  
55 Their ability to reproduce the statistics of high dimensional models of such  
56 quantities as probability density and autocorrelation functions has been shown  
57 by Franzke et al. (2005) and Franzke and Majda (2006). These systematic  
58 reduced order models are also skillful in reproducing the extreme value statis-  
59 tics and the predictability of extreme events of higher dimensional systems  
60 (Franzke, 2012).

61 The idea behind tipping point prediction is that the underlying essential  
62 dynamics can be represented by a potential well driven by additive white  
63 noise (Scheffer et al., 2009; Livina et al., 2013). This is based on bifurca-  
64 tion theory of low-dimensional Ordinary Differential Equations. However,  
65 in practical situations one has often an one-dimensional indicator time se-  
66 ries (e.g. data from an ice core, measurement of the Meridional Overturning  
67 Circulation in the ocean) of a complex high-dimensional system. Using such  
68 a time series for tipping point prediction implicitly assumes either a weak  
69 coupling or time scale separation between the indicator time series and the  
70 remaining variables of the system. In this approach there is also a hidden as-  
71 sumption of an additive coupling between the observed indicator time series  
72 and the rest of the system. These approaches do not consider the possibility  
73 of multiplicative coupling which could lead to a state-dependent noise. For  
74 instance, this state-dependent noise could create a double well potential on  
75 its own with the deterministic dynamics playing no role in the creation of  
76 the double well potential (Sura et al., 2005) (see their figure 1). In such a  
77 system all transitions are purely noise driven. This illustrates the danger on  
78 relying upon purely data driven approaches since they are unlikely to be able  
79 to distinguish the dynamical causes of the potential well. The here proposed  
80 approach of using dynamical models also provides insight into the underlying  
81 dynamics and thus gives more confidence in the predictions.

82 In this study we will derive dynamical reduced order models which are  
 83 driven by a slow forcing towards a bifurcation point. After demonstrating  
 84 that the reduced order model reproduce the same tipping point behaviour we  
 85 will elucidate the roles of time scale separation and coupling strength between  
 86 resolved and unresolved modes and how they affect the tipping behaviour:  
 87 whether the systems undergoes a clean tipping or flickers between the two  
 88 equilibrium states. Flickering has recently also been proposed as an indicator  
 89 of an imminent tipping event (Veraart et al., 2012). Both kinds of behaviour  
 90 have been found to precede critical transitions. So far no explanation has  
 91 been given on which properties of the underlying dynamics they depend.  
 92 To elucidate the conditions under which one can expect a clean tipping or  
 93 flickering is a major motivation of this study.

94 We will introduce the stochastic conceptual model which represents a  
 95 minimal prototype climate model in Section 2. We discuss its performance  
 96 when driven by time-dependent forcing in Section 3 and its ability to robustly  
 97 predict tipping points in Section 4. In Section 4 we also discuss the robust-  
 98 ness of the typically used tipping point prediction methods. We provide a  
 99 summary of our results in section 5.

## 100 2. Stochastic Conceptual Model

101 In this section we describe the conceptual model which we are using in  
 102 our study of tipping points. A similar version of this conceptual model has  
 103 been used in previous studies (Majda et al., 2005, 2008; Franzke et al., 2007;  
 104 Franzke, 2012). The conceptual model is 4 dimensional and contains the  
 105 essential dynamics of more complex climate models even though it is of much  
 106 lower dimensionality.

107 The conceptual model we are using in this study has two slow or cli-  
 108 mate variables denoted by  $(x_1, x_2)$ . These two modes evolve slower than  
 109 the other two modes  $(y_1, y_2)$ . These two fast modes represent turbulent ed-  
 110 dies and convective systems in the climate system which are in many climate  
 111 models not fully resolved. In realistic systems there would be innumerable  
 112 many fast modes, and in order to mimic their combined effect on the two  
 113 slow climate modes we include damping and stochastic forcing  $-\frac{\gamma}{\varepsilon}y + \frac{\sigma}{\sqrt{\varepsilon}}dW$   
 114 in the equations for  $y$  where  $W$  denotes a Wiener process. This approxi-  
 115 mation is motivated by the fact that these fast modes are associated with  
 116 turbulent energy transfers and strong mixing. In this study we do not re-  
 117 quire a detailed description of these processes because we are only interested

118 in their combined effect on the slow resolved modes and not in their detailed  
 119 evolution. The stochastic climate model is given by

$$dx_1 = ((-x_2(L_{12} + a_1x_1 + a_2x_2) + d_1x_1 + F_1(t)) \quad (1a)$$

$$+ \theta(L_{13}y_1 + b_{123}x_2y_1 + (c_{131} + c_{113})x_1y_1)) dt \quad (1b)$$

$$dx_2 = ((+x_1(L_{21} + a_1x_1 + a_2x_2) + d_2x_2 + F_2(t)) \quad (1c)$$

$$+ \theta(L_{24}y_2 + b_{213}x_1y_1 + (e_{242} + e_{224})x_2y_2)) dt \quad (1d)$$

$$dy_1 = \left(-L_{13}x_1 + b_{312}x_1x_2 + c_{311}x_1x_1 + F_3(t) - \frac{\gamma_1}{\varepsilon}y_1\right) dt + \frac{\sigma_1}{\sqrt{\varepsilon}}dW_1 \quad (1e)$$

$$dy_2 = \left(-L_{24}x_2 + e_{422}x_2x_2 + F_4(t) - \frac{\gamma_2}{\varepsilon}y_2\right) dt + \frac{\sigma_2}{\sqrt{\varepsilon}}dW_2 \quad (1f)$$

120 The parameter  $\varepsilon$  controls the time-scale separation between the slow and  
 121 fast variables. Energy conservation of the nonlinear operator requires that  
 122  $b_{123} + b_{213} + b_{312} = 0$ ,  $c_{131} + c_{113} + c_{311} = 0$  and  $e_{242} + e_{224} + e_{422} = 0$ . The linear  
 123 operator matrix  $\mathbf{L}$  is skew-symmetric. The climate and fast modes are both  
 124 linearly and nonlinearly coupled through triad and dyad interactions. Note  
 125 that the forcing  $\mathbf{F}(t)$  is time-dependent in contrast to earlier studies (Majda  
 126 et al., 2005, 2008; Franzke et al., 2007; Franzke, 2012). We added a parameter  
 127  $\theta$  in (1) with which we are able to control the strength of the interaction  
 128 between the deterministic nonlinear dynamics and the fast unresolved modes.

129 To highlight the structural form of our conceptual model (1) we rewrite  
 130 it as

$$d\mathbf{z} = (\mathbf{F}(t) + L\mathbf{z}(t) + B(\mathbf{z}(t), \mathbf{z}(t))) dt + \sigma d\mathbf{W}. \quad (2)$$

131 This is the same structural form as climate models have with a forcing  $\mathbf{F}$ ,  
 132 a linear operator  $L$ , a quadratic nonlinear operator  $B$  and additive noise  
 133 forcing  $d\mathbf{W}$ . While most current climate models are deterministic there are  
 134 a few numerical weather prediction models which have stochastic terms.

### 135 2.1. Explicit Stochastic Mode Reduction

136 We now apply the systematic stochastic mode reduction procedure (Ma-  
 137 jda et al., 1999, 2001, 2002) to the model (1) to obtain explicit reduced  
 138 stochastic equations for the slow variables  $\mathbf{x}$ . The simplicity of the above  
 139 model allows us to do the stochastic mode reduction directly using the equa-  
 140 tions without transforming it to the corresponding Fokker-Planck equation.

141 The stochastic differential equation (SDE) for the variable  $\mathbf{y}$  in (1) is  
 142 linear in  $\mathbf{y}$ . Thus, given  $\mathbf{x}(t)$  its solution is

$$y_1(t) = e^{-\frac{\gamma_1 t}{\varepsilon}} y_1(0) + \int_0^t e^{-\frac{\gamma_1(t-s)}{\varepsilon}} [-L_{13}x_1(s) + b_{312}x_1(s)x_2(s) + c_{311}x_1(s)x_1(s) + F_3(s)] ds + g_1(t) \quad (3)$$

143 where

$$g_1(t) = \frac{\sigma_1}{\sqrt{\varepsilon}} \int_0^t e^{-\frac{\gamma_1(t-s)}{\varepsilon}} dW_1(s) \quad (4)$$

144 and

$$y_2(t) = e^{-\frac{\gamma_2 t}{\varepsilon}} y_2(0) + \int_0^t e^{-\frac{\gamma_2(t-s)}{\varepsilon}} [-L_{24}x_2(s) + e_{422}x_2(s)x_2(s) + F_4(s)] ds + g_2(t) \quad (5)$$

145 where

$$g_2(t) = \frac{\sigma_2}{\sqrt{\varepsilon}} \int_0^t e^{-\frac{\gamma_2(t-s)}{\varepsilon}} dW_2(s) \quad (6)$$

146 Inserting (3) and (5) into the first two equations in (1) for the variable  $\mathbf{x}$   
 147 yields an exact, non-Markovian system of equations for  $\mathbf{x}(t)$ .

148 Since we are interested in the long time statistical behaviour of the climate  
 149 variables  $\mathbf{x}(t)$  as  $\varepsilon \rightarrow 0$ , we consider the asymptotic limit as  $\varepsilon \rightarrow 0$  of the  
 150 three terms on the right of (3) and (5). First we immediately have

$$e^{-\frac{\gamma_1 t}{\varepsilon}} y_1(0) \rightarrow 0 \quad (7)$$

$$e^{-\frac{\gamma_2 t}{\varepsilon}} y_2(0) \rightarrow 0 \quad (8)$$

151 Second, using integration by parts we find

$$\begin{aligned} & \int_0^t e^{-\frac{\gamma_1(t-s)}{\varepsilon}} [-L_{13}x_1(s) + b_{312}x_1(s)x_2(s) + c_{311}x_1(s)x_1(s) + F_3(s)] ds \\ & \rightarrow \frac{\varepsilon}{\gamma_1} [-L_{13}x_1(t) + b_{312}x_1(t)x_2(t) + c_{311}x_1(t)x_1(t) + F_3(t)] \quad (9) \end{aligned}$$

$$\begin{aligned} & \int_0^t e^{-\frac{\gamma_2(t-s)}{\varepsilon}} [-L_{24}x_2(s) + e_{422}x_2(s)x_2(s) + F_4(s)] ds \\ & \rightarrow \frac{\varepsilon}{\gamma_2} [-L_{24}x_2(t) + e_{422}x_2(t)x_2(t) + F_4(t)] \quad (10) \end{aligned}$$



152 Finally, it can be shown that  $g_1(t)$  and  $g_2(t)$  are itself approximately white  
 153 noise as  $\varepsilon \rightarrow 0$  (Majda et al., 2001)

$$g_1(t)dt \rightarrow \sqrt{\varepsilon} \frac{\sigma_1}{\gamma_1} dW_1(t) \quad (11)$$

$$g_2(t)dt \rightarrow \sqrt{\varepsilon} \frac{\sigma_2}{\gamma_2} dW_2(t) \quad (12)$$

154 for this we use the fact that  $g_1(t)$  and  $g_2(t)$  are Gaussian and the two prop-  
 155 erties for any test function  $\eta$

$$\mathbb{E} \left( \frac{1}{\varepsilon} \int_0^\infty \eta(t) g_j(t) dt \right) = 0 \quad (13)$$

156 and

$$\mathbb{E} \left( \frac{1}{\varepsilon} \int_0^\infty \eta(t) g_i(t) dt \right) \left( \frac{1}{\varepsilon} \int_0^\infty \eta(t) g_j(t) dt \right) \rightarrow \varepsilon \frac{\sigma_j^2}{\gamma_j^2} \delta_{ij} \int_0^\infty \eta^2(t) dt \quad (14)$$

157 We note, however, that, as an approximation of a process with finite corre-  
 158 lation time,  $dW_i(t)$  has to be interpreted in the Stratonovich sense (Gardiner,  
 159 1985).

160 Combining these formulas in the first two equations of (1), we obtain the  
 161 following SDE transformed to Itô form with the noise induced drift (Gardiner,

162 1985):

$$\begin{aligned}
dx_1(t) = & (-x_2(t) (L_{12} + a_1x_1(t) + a_2x_2(t)) + d_1x_1(t) + F_1(t)) dt \\
& + \theta \left( \frac{\varepsilon}{\gamma_1} (L_{13}F_3(t) - L_{13}L_{13}x_1(t) + b_{123}F_3(t)x_2(t) \right. \\
& + L_{13}b_{312}x_1(t)x_2(t) - L_{13}b_{123}x_1(t)x_2(t) + L_{13}c_{311}x_1^2(t) \\
& + b_{312}b_{123}x_1(t)x_2^2(t) + b_{123}c_{311}x_2(t)x_1^2(t) + (c_{131} + c_{113}) \\
& (c_{311}x_1^3(t) - L_{13}x_1^2(t) + b_{312}x_1^2(t)x_2(t) + F_3(t)x_1(t)) \left. \right) dt \\
& + \varepsilon \frac{1}{2} \frac{\sigma_1^2}{\gamma_1^2} (b_{213}b_{123}x_1(t) + (L_{13} + b_{123}x_2(t) \\
& + (c_{131} + c_{113}) x_1(t)) (c_{131} + c_{113}) \left. \right) dt \\
& + \sqrt{\varepsilon} \frac{\sigma_1}{\gamma_1} (L_{13} + b_{123}x_2(t) + (c_{131} + c_{113}) x_1(t)) dW_1(t) \quad (15a)
\end{aligned}$$

$$\begin{aligned}
dx_2(t) = & (x_1(t) (L_{21} + a_1x_1(t) + a_2x_2(t)) + d_2x_2(t) + F_2(t)) dt \\
& + \theta \left( \frac{\varepsilon}{\gamma_2} (L_{24}F_4(t) - L_{24}L_{24}x_2(t) + L_{24}e_{422}x_2^2(t) \right. \\
& + (e_{242} + e_{224}) (e_{422}x_2^3(t) - L_{24}x_2^2(t) + F_4(t)x_2(t)) \left. \right) dt \\
& + \frac{\varepsilon}{\gamma_1} (-b_{213}L_{13}x_1(t)x_1(t) + b_{213}c_{311}x_1^3(t) \\
& + b_{213}b_{312}x_1(t)x_1(t)x_2(t) + b_{213}F_3(t)x_1(t)) dt \\
& + \varepsilon \frac{1}{2} \frac{\sigma_1^2}{\gamma_1^2} (b_{213}b_{123}x_2(t) + L_{13}b_{213} + (c_{131} + c_{113}) b_{213}x_1(t)) dt \\
& + \varepsilon \frac{1}{2} \frac{\sigma_2^2}{\gamma_2^2} (L_{24} + (e_{242} + e_{224}) x_2(t)) (e_{242} + e_{224}) dt \\
& + \sqrt{\varepsilon} \frac{\sigma_1}{\gamma_1} b_{213}x_1(t) dW_1(t) \\
& + \sqrt{\varepsilon} \frac{\sigma_2}{\gamma_2} (L_{24} + (e_{242} + e_{224}) x_2(t)) dW_2(t) \quad (15b)
\end{aligned}$$

163 Note that coarse graining time as  $t \rightarrow \frac{t}{\varepsilon}$  amounts to setting  $\varepsilon = 1$  (Majda  
164 et al., 1999, 2001; Franzke et al., 2005; Franzke and Majda, 2006).

165 To highlight the structural differences we rewrite the reduced model in  
166 the following form

$$\frac{d\mathbf{x}}{dt} = \tilde{\mathbf{F}}(t) + \tilde{L}\mathbf{x}(t) + \tilde{B}(\mathbf{x}(t), \mathbf{x}(t)) + \tilde{M}(\mathbf{x}(t), \mathbf{x}(t), \mathbf{x}(t)) + \sigma_1 d\mathbf{W}_1(t) + \sigma_2(\mathbf{x}(t)) d\mathbf{W}_2(t). \quad (16)$$

167 The qualitative new terms are the deterministic cubic operator  $\tilde{M}$  and the  
 168 state-dependent noise  $\sigma_2(\mathbf{x})$ . In general the deterministic cubic operator acts  
 169 as effective damping while it also allows the system to be linearly unstable  
 170 (Majda et al., 2009).

### 171 2.2. Nonlinear Deterministic Dynamics

172 The nonlinear deterministic dynamics for the climate variables  $x_1$  and  $x_2$   
 173 of the conceptual climate model (1) is given by

$$\frac{dx_1}{dt} = -x_2(L_{12} + a_1x_1 + a_2x_2)dt + d_1x_1dt + F_1dt \quad (17a)$$

$$\frac{dx_2}{dt} = x_1(L_{21} + a_1x_1 + a_2x_2)dt + d_2x_2dt + F_2dt. \quad (17b)$$

174 Here we set  $y_1 = y_2 = 0$  in order to explore the bifurcation behaviour of the  
 175 climate modes.

176 By varying the forcing  $\mathbf{F}$  the system (15) undergoes several bifurcations  
 177 as shown by Majda et al. (2005). By decreasing  $F_1$  starting from -0.1 the  
 178 system undergoes first a saddle-node bifurcation with two stable states, at  
 179 the second bifurcation point a homoclinic orbit appears before it undergoes a  
 180 supercritical Hopf bifurcation. For more details about the bifurcation struc-  
 181 ture of the nonlinear deterministic dynamics (17) see Majda et al. (2005)  
 182 (See their Fig. 3.2 for the bifurcation diagram).

### 183 2.3. Model Integration Details

184 A 5000 member ensemble is created by integrating the model for  $10^5$   
 185 time units starting from 5000 different initial conditions which were chosen  
 186 randomly and using different stochastic noise realizations. To integrate the  
 187 model we are using a fourth order Runge-Kutta scheme for the deterministic  
 188 part and an Euler forward scheme for the stochastic part. We use a time  
 189 step of  $10^{-4}$  time units and save output every  $\frac{1}{8}$  time unit.

190 Furthermore, we use the three time scale separation values  $\varepsilon = 0.1, 0.5$   
 191 and 1.0. These three cases correspond to time scale separation ( $\varepsilon = 0.1$ ),  
 192 moderate time scale separation ( $\varepsilon = 0.5$ ) and no time scale separation  
 193 ( $\varepsilon = 1.0$ ). The reduced order model is only valid in case of time scale sep-  
 194 aration but in most natural system we have only moderate or no time scale  
 195 separation. For instance, the atmospheric circulation has  $\varepsilon$  values between  
 196 0.6 and 1.0 (e.g. Franzke et al. (2005); Franzke and Majda (2006)). Thus,  
 197 we also have to test how well the method works in these more realistic cases.

198 A series of studies has shown that the stochastic mode reduction performs  
199 reasonably well also for moderate or no time scale separation (Majda et al.,  
200 2002, 2005, 2008; Franzke et al., 2005; Franzke and Majda, 2006; Franzke,  
201 2012).

### 202 3. Prediction of non-Stationary Dynamics

203 In this section we will evaluate how well the reduced stochastic models  
204 reproduce the full dynamics when driven by a time-dependent forcing  $\mathbf{F}(t)$ .  
205 First we use  $F_1(t) = -0.2 + 0.4 * \sin(t/5000)$  which is a periodic forcing  
206 on a very slow time scale. The forcing has been chosen in such a way that  
207 it passes through all bifurcation points. Typical realizations can be seen  
208 in Fig. 1 for three different time scale separations. As can be seen the  
209 conceptual model exhibits different dynamical regimes for different forcing  
210 values. Furthermore, the reduced model captures this behavior very well  
211 for all three time scale separation values  $\varepsilon$ . This is further confirmed by the  
212 autocorrelation function and the PDF. The reduced order model captures the  
213 decay of the autocorrelation function (Fig. 2) and of the PDFs (Fig. 3) very  
214 well for all time scale separations. For  $\varepsilon = 0.1$  the full and reduced dynamics  
215 are almost indistinguishable. The shape of the highly non-Gaussian PDFs  
216 are very well captured for both the marginal and joint PDFs by the reduced  
217 model (Fig. 4) for all three values of  $\varepsilon$ . Also in the case with no time scale  
218 separation  $\varepsilon = 1.0$  the PDF is very well captured which is a very promising  
219 result since in realistic systems one rarely has time scale separation.

220 Now we evaluate how well the reduced model performs if the time periodic  
221 forcing drives one of the fast modes. We use  $F_3(t) = -0.2 + 0.4 * \sin(t/5000)$   
222 which is a slow periodic driving of the fast mode  $y_1$ . Also in this case the  
223 reduced stochastic model reproduces the statistics of the full dynamics very  
224 well. Again the autocorrelation function is extremely well captured for  $\varepsilon =$   
225  $0.1$  and still well for  $\varepsilon = 0.5$  and  $\varepsilon = 1.0$  (Fig. 5). To put this into context, the  
226 stochastic mode reduction strategy is strictly valid only in the limit  $\varepsilon \rightarrow 0$  but  
227 as our empirical results show it still performs well in cases with no time scale  
228 separation at all. This is a promising result suggesting that our proposed  
229 approach will also work for observed data which likely has only moderate  
230 time scale separation.

#### 231 4. Prediction of Tipping Points

232 In this section we discuss the role of time scale separation and coupling  
 233 strength and how well the reduced order models predict tipping points by  
 234 driving the models with a linearly increasing forcing  $F_1(t) = -0.5 + 0.00002 * t$ .  
 235 In figure 6 we display two example trajectories one for weak ( $\theta = 0.1$ ) and one  
 236 for strong ( $\theta = 1.0$ ) coupling between climate and fast modes. In both cases  
 237 we set  $\varepsilon = 0.1$ . The here relevant major difference between the two cases is  
 238 the level of variability; for weak coupling the variability is much smaller and  
 239 the system tips later. Furthermore, the reduced dynamics capture the full  
 240 dynamics again very well.

241 Another way of looking at the non-stationary evolution of the conceptual  
 242 model is to compute time evolving PDFs. Here we compute the PDF at a  
 243 fixed time  $t$  over the 5000 member ensemble. This will reveal how narrow  
 244 the window of tipping is and how well the reduced dynamics captures this  
 245 essential part of the non-stationary behavior.

246 A comparison of the time evolving marginal PDFs (Figs. 7 and 8) shows  
 247 that for weak coupling there is a rather sharply defined tipping time because  
 248 the PDFs are very narrow and do not overlap during the state transition.  
 249 This is also very well captured by the reduced dynamics. This is different  
 250 in the case of strong coupling (Fig. 8) where the PDF is much broader and  
 251 there is a rather smooth transition between the two states indicating that the  
 252 tipping time is not well defined and the system tends to tip early or jumps a  
 253 couple of times between both equilibrium states before it settles down on the  
 254 surviving equilibrium state. However, in the case of time scale separation  
 255 ( $\varepsilon = 0.1$ ) the PDF is much sharper, though not as sharp as for the weak  
 256 coupling case. Also the time window when the system tips is much narrower  
 257 than for moderate time scale separation. This shows that the typical tipping  
 258 point prediction methods are likely to only robustly work in the case of time  
 259 scale separation and weak coupling. In other situations, which are likely  
 260 more realistic, the system might not undergo a clear critical transition and  
 261 flickers between the two equilibrium states.

262 The reduced dynamics reproduces the full dynamics very well in the case  
 263 of time scale separation ( $\varepsilon = 0.1$ ) and reasonably well in the other two cases  
 264 for both weak and strong coupling. This shows the reduced order models can  
 265 play a useful role in predicting tipping points.

266 The traditional tipping point indicators are sensitive to the way the time  
 267 series is detrended and the length of the window length. Here we can use

268 the ensemble to test whether ensemble averaging detects the critical slowing  
 269 down signal. For this purpose we compute the lag-1 correlation coefficient  
 270  $\langle x(t)x(t-1) \rangle$  as used in the AR(1) tipping indicators and the variance  
 271  $\langle x(t)^2 \rangle$  at time  $t$ , where  $\langle \rangle$  denotes an ensemble average. This approach  
 272 has the advantage that no detrending is necessary, since any trend between  
 273 two consecutive time points will be negligibly small, and we also do not have  
 274 to define a window length for the averaging.

275 Fig. 9 shows one time series realization for both climate modes in the  
 276 case of weak coupling which we consider to be the truth here together with  
 277 the ensemble lag-1 correlation and variance tipping point indicators averaged  
 278 over 1000 ensemble members. The variance indicator increases in magnitude  
 279 when approaching the tipping time. The amount of time scale separation  
 280 determines when the variance reaches its maximum. For  $\varepsilon = 0.1$  the variance  
 281 reaches its maximum at about the time of tipping while for  $\varepsilon = 1.0$  it reaches  
 282 its maximum before the time of tipping and actually already decreases before  
 283 the tipping. On the other hand, the lag-1 indicator increases in value only  
 284 for  $\varepsilon = 0.1$ . In this case it also reaches its maximum before the tipping  
 285 time and starts the decrease by the time of tipping. For the other two  
 286 cases the lag-1 indicator is rather flat or only minimally increasing. Thus,  
 287 our model results suggest that the amount of time scale separation and the  
 288 coupling strength between resolved and unresolved modes determine whether  
 289 the critical transition is due to critical slowing down or flickering.

290 In the case of strong coupling both indicators increase before the time  
 291 of tipping in the case of  $\varepsilon = 0.1$  whereas in the other two cases, with only  
 292 moderate or no time scale separation, the system jumps a few times between  
 293 both equilibrium states (Fig. 10). But in these two cases both indicators  
 294 seem to peak at about the time that the start of the transition to the other  
 295 equilibrium point becomes visible in the PDF (Fig. 8). This suggests that  
 296 the ensemble approach can still be useful in predicting the onset of the switch  
 297 to another equilibrium point even though it is not a clear tipping point but  
 298 flickering. This suggests that the ensemble approach might be useful as an  
 299 early warning system even though there will be no clear or unique time of  
 300 tipping.

301 Now we discuss how well the traditional tipping point indicators perform.  
 302 In Fig. 11 we display the results of the full dynamics simulations from us-  
 303 ing 4 typical tipping point indicators: AR(1), variance, skewness and linear  
 304 decay rate derived from the quasi-stationary density (Gardiner, 1985; Livina  
 305 et al., 2012; Sieber and Thompson, 2012). To compute these indicators we

306 use a sliding window of length 1000 and then linearly detrend the time series  
 307 in each window. Here we average over the window length and not the en-  
 308 semble. While the AR(1), variance and skewness are standard quantities the  
 309 quasi-stationary density involves the Fokker-Planck equation. This indicator  
 310 assumes that the deterministic dynamics of the detrended time series evolves  
 311 in a potential well  $U(x)$  (Gardiner, 1985; Sieber and Thompson, 2012). The  
 312 linear decay rate  $\kappa$  can then be computed via

$$\frac{1}{2}\partial_x p(x) = -\kappa x p(x) + c \quad (18)$$

313 where  $p(x)$  denotes the empirical density and  $c$  a constant. We approximate  
 314 the derivative of the density with finite differences. We apply these 4 indi-  
 315 cators to 100 ensemble time series for the weak coupling and  $\varepsilon = 0.1$  case  
 316 which can be considered to be the best case scenario for tipping point pre-  
 317 dictions. During the displayed time range in Fig. 11 both stable equilibria  
 318 exist (compare with Fig. 6).

319 The results display a wide variety of tipping indicator behavior. It is  
 320 clearly visible that, even though the PDFs (Fig. 8c) show a very narrow tip-  
 321 ping time range, the indicators do not seem to robustly signal the imminent  
 322 tipping point in our model experiments. For some realizations the indicators  
 323 do not predict a tipping at all while when they predict a tipping the timing  
 324 varies widely (Fig. 11). This is the case for all 4 tipping point indicators.  
 325 At least for this model experiment our proposed ensemble model prediction  
 326 system seems to perform more robustly and reliably.

## 327 5. Summary

328 Using a conceptual model mimicing aspects of complex climate mod-  
 329 els we elucidated the tipping point behaviour and how it depends on time  
 330 scale separation and the coupling strength between resolved and unresolved  
 331 modes. We find that for model experiments the theory of critical slowing  
 332 down applies best to the case of large time scale separation and weak cou-  
 333 pling between resolved and unresolved modes. In this situation there is a  
 334 clear and distinct tipping event. For moderate or small time scale separation  
 335 and strong coupling the model flickers between the two equilibrium states.  
 336 Both critical slowing down (Scheffer et al., 2009; Sieber and Thompson, 2012;  
 337 Livina et al., 2012) and flickering (Scheffer et al., 2009; Lenton, 2011) have  
 338 been proposed as indicators of imminent tipplings and here we have shown

339 which properties of the underlying dynamics are responsible for these two  
340 distinct behaviours preceding a critical transition.

341 Furthermore, we have shown that reduced order models are able to repro-  
342 duce the tipping point behaviour of more complex models. Our model results  
343 suggest that predicting the time of tipping works best for systems with time  
344 scale separation and weak coupling between the resolved and the unresolved  
345 part of the system. As can be seen in Fig. 10 for strong coupling and lack of  
346 time scale separation the system flickers between the two equilibrium states.  
347 The reduced order models well reproduce this flickering.

348 A potential advantage of the proposed dynamical tipping prediction ap-  
349 proach is that the reduced order models can be run in forecast mode with  
350 extrapolation of the forcing. These ensemble predictions will provide a prob-  
351 abilistic forecast of the tipping time which can then be used in integrated  
352 assessment and decision making models. This is not possible with the diag-  
353 nostic tipping indicators which cannot provide any estimate of the tipping  
354 time other than that the system might approach the tipping point. Further-  
355 more, the extrapolation can be done also with an ensemble of possible and  
356 plausible forcings or control parameters. This ensemble can then be used to  
357 make probabilistic forecasts about whether a tipping is imminent or not.

358 The stochastic mode reduction approach described here requires the knowl-  
359 edge of the dynamical equations of the system of interest. For the climate  
360 system the normal forms of stochastic climate models have been derived by  
361 Majda et al. (2009). In order to estimate the necessary parameter values  
362 of the stochastic differential equation from data one can use Bayesian infer-  
363 ence methods which also takes proper account of all uncertainties (Peavoy  
364 et al., 2013). However, these approaches need to be extended to work in a  
365 non-stationary setting. Conceptually this is straight forward by treating the  
366 forcing or the control parameter as an additional equation. This forcing or  
367 the control parameter are not necessarily directly observable. In this  
368 case it can be treated as a latent variable in Bayesian inference (Peavoy et  
369 al., 2013). This dynamical model fitting approach also offers the possibility  
370 of further insight into the underlying mechanisms of observed critical tran-  
371 sitions. For instance, the structure of the additional equation describing the  
372 forcing could either be an increasing or decreasing function or a noise driven  
373 stationary process. In the latter case all critical transitions would likely be  
374 noise induced.

375 In many areas of science the evolution equations are not known. In this  
376 situation one can use non-parametric approaches to estimate the evolution



377 equations just from observed data (Crommelin and Vanden-Eijnden, 2006;  
378 Carpenter and Brock, 2011) or fit a potential well type equation with ad-  
379 ditive noise (Livina et al., 2010, 2011; Sieber and Thompson, 2012). These  
380 approaches are more general and can be applied to many observational data  
381 sets.

382 Our results suggest that any early warning system of tipping points should  
383 include an ensemble approach using dynamical models. This would allow for  
384 a probabilistic prediction of imminent tipping points and would provide an  
385 estimated range of tipping times which might be useful to decide on the best  
386 avoidance or mitigation strategies by taking all uncertainties into account.

387 **Acknowledgments** I thank Drs. V. Livina and P. Ditlevsen for their  
388 comments on an earlier draft of this manuscript which helped to improve it.  
389 This study is part of the British Antarctic Survey Polar Science for Planet  
390 Earth Programme. It was funded by The Natural Environment Research  
391 Council and through NERC grant NE/G015317/1.

392 **References**

393 **References**

- 394 R. Biggs, S. R. Carpenter and W. A. Brock, Turning back from the brink:  
395 Detecting an impending regime shift in time to avert it. *Proc. Nat. Acad.*  
396 *Sci. USA*, 106 (2009), 826-831.
- 397 C. Boettinger and A. Hastings, Early warning signals and the prosecutor's  
398 fallacy. *Proc. Roy. Soc. B*, (2012a), doi:10.1098/rspb.2012.2085.
- 399 C. Boettinger and A. Hastings, Quantifying limits to detection of  
400 early warning for critical transitions. *J. R. Soc. Interface*, (2012b),  
401 doi:10.1098/rspb.2012.2085.
- 402 S. R. Carpenter and W. A. Brock, Early warnings of unknown nonlinear  
403 shifts: a nonparametric approach. *Ecology*, 92 (2011), 2196-2201.
- 404 A. A. Cimatoribus, S. S. Drixfhout, V. Livina and G. van der Schrier,  
405 Dansgaard-Oeschger events: bifurcation points in the climate system.  
406 *Clim. Past.*, 9 (2013), 323-333.
- 407 D. Crommelin and E. Vanden-Eijnden, Reconstruction of diffusions using  
408 spectral data from timeseries. *Comm. Math. Sci.*, 4 (2006), 651-668.
- 409 V. Dakos, M. Scheffer, E. H. van Nes, V. Brovkin, V. Petoukhov and H.  
410 Held, Slowing down as an early warning signal for abrupt climate change.  
411 *Proc. Nat. Acad. Sci. USA*, 105, 14308-14312 (2008)
- 412 P. D. Ditlevsen and S. J. Johnsen, Tipping points: Early warning and wishful  
413 thinking. *Geophys. Res. Lett.*, 37, L19703 (2010)
- 414 I. Eisenmann and J. S. Wettlaufer, Nonlinear threshold behavior during the  
415 loss of Arctic sea ice. *Proc. Nat. Acad. Sci. USA*, 106 (2008), 28-32.
- 416 C. Franzke, A. J. Majda and E. Vanden-Eijnden, 2005: Low-Order Stochastic  
417 Mode Reduction for a Realistic Barotropic Model Climate. *J. Atmos. Sci.*,  
418 62, 1722-1745.
- 419 C. Franzke and A. J. Majda, 2006: Low-Order Stochastic Mode Reduction  
420 for a Prototype Atmospheric GCM. *J. Atmos. Sci.*, 63, 457-479.

- 421 C. Franzke, A. J. Majda and G. Branstator, 2007: The origin of nonlinear  
422 signatures of Planetary Wave Dynamics: Mean Phase Space Tendencies  
423 and Contributions from Non-Gaussianity. *J. Atmos. Sci.*, 64, 3987-4003.  
424 doi: 10.1175/2006JAS2221.1
- 425 C. Franzke, Predictability of extreme events in a nonlinear stochastic-  
426 dynamical model. *Phys. Rev. E*, 85 (2012) doi: 10.1103/Phys-  
427 RevE.85.031134
- 428 C. W. Gardiner, Handbook of stochastic methods, 1985. Springer-Verlag,  
429 442pp.
- 430 V. Guttal and C. Jayaprakash, Changing skewness: an early warning signal  
431 of regime shifts in ecosystems. *Ecology Lett.*, 11 (2009), 450-460.
- 432 H. Held and T. Kleinen, Detection of climate system bifurcations by de-  
433 generate fingerprinting. *Geophys. Res. Lett.*, 31, L23207 (2004) doi:  
434 10.1029/2004GL020972
- 435 C. Kuehn, A mathematical framework for critical transitions: Bifurcations,  
436 fast-slow systems and stochastic dynamics. *Physica D*, 240 (2011) 1020-  
437 1035.
- 438 C. Kuehn, A mathematical framework for critical transitions: Normal forms,  
439 variance and applications. *J. Nonlin. Sci.*, 23 (2013) 457-510.
- 440 T. M. Lenton, H. Held, E. Kriegler, J. W. Hall, W. Lucht, S. Rahmstorff and  
441 H. J. Schellnhuber, Tipping elements in the Earth's climate system. *Proc.*  
442 *Nat. Acad. Sci. USA*, 105 (2008) 1786-1793
- 443 T. M. Lenton, Early warning of climate tipping points. *Nature Climate*  
444 *Change*, 1 (2011) 201-209
- 445 V. N. Livina and T. M. Lenton, A modified method for detecting incipient  
446 bifurcations in a dynamical system. *Geophys. Res. Lett.*, L03712 (2007)  
447 doi: 10.1029/2006GL028672
- 448 V. N. Livina, F. Kwasniok and T. M. Lenton, Potential analysis reveals  
449 changing number of climate states during the last 60kyr. *Clim. Past*, 6,  
450 77-82. (2010).

- 451 V. N. Livina, F. Kwasniok, G. Lohmann, J. W. Kantelhardt and T. M.  
452 Lenton, Changing climate states and stability: from pliocene to present,  
453 *Clim. Dyn.*, 37, 2437-2453. (2011).
- 454 V. N. Livina, P. D. Ditlevsen and T. M. Lenton, An independent test of  
455 methods of detecting system states and bifurcations in time-series data.  
456 *Physica A*, 391, 485-496. (2012).
- 457 V. N. Livina, G. Lohmann, M. Muddelsee and T. M. Lenton, Forecasting  
458 the underlying potential governing the time series of a dynamical system.  
459 *Physica A*, in press. doi: 10.1016/j.physa.2013.04.036
- 460 A. J. Majda, I. Timofeyev and E. Vanden-Eijnden, 1999: Models for stochas-  
461 tic climate prediction. *Proc. Nat. Acad. Sci. USA*, 96, 14687-14691.
- 462 A. J. Majda, I. Timofeyev and E. Vanden-Eijnden, 2001: A mathematical  
463 framework for stochastic climate models. *Commun. Pure Appl. Math.*, 54,  
464 891-974.
- 465 A. J. Majda, I. Timofeyev and E. Vanden-Eijnden, 2002: A priori tests of a  
466 stochastic mode reduction strategy. *Physica D*, 170, 206-252.
- 467 A. J. Majda, R. Abramov and M. Grote, 2005: *Information Theory and*  
468 *Stochastics for Multiscale Nonlinear Systems*. CRM Monograph Series,  
469 American Mathematical Society.
- 470 A. J. Majda, C. Franzke and B. Khouider, 2008: An applied mathematics  
471 perspective on stochastic modelling for climate. *Phil. Trans. R. Soc. A*,  
472 366, 2429-2455.
- 473 A. J. Majda, C. Franzke and D. Crommelin, 2009: Normal forms for reduced  
474 stochastic climate models. *Proc. Nat. Acad. Sci. USA*, 106, 3649-3653. doi:  
475 10.1073/pnas.0900173106
- 476 P. E. McSharry, L. A. Smith and L. Tarassenko: Prediction of epileptic  
477 seizures: are nonlinear methods relevant? *Proc. Nature Med.* (2003), 9,  
478 241-242.
- 479 A. H. Monahan, J. Alexander and A. J. Weaver: Stochastic models of the  
480 meridional overturning circulation: time scales and patterns of variability.  
481 *Phil. Trans. Roy. Soc. A* (2008), 366, 2527-2544.

- 482 D. Peavoy, C. Franzke and G. Roberts: Systematic Physics Constrained  
483 Parameter Estimation of Stochastic Differential Equations. *Comp. Stat.*  
484 *Data Anal.* (2013), in preparation.
- 485 J. Ridley, J. M. Gregory, P. Huybrechts and J. Lowe: Thresholds for irre-  
486 versible decline of the Greenland ice sheet. *Clim. Dyn.* (2010), 35, 1049-  
487 1057. doi: 10.1007/s00382-0646-0
- 488 M. Scheffer, S. R. Carpenter, J. A. Foley, C. Folke and B. Walker, Catastrophic  
489 shifts in ecosystems. *Nature*, 413 (2001), 591-596.
- 490 M. Scheffer, J. Bascompte, W. A. Brock, V. Brovkin, S. R. Carpenter, V.  
491 Dakos, H. Held, E. H van Mes, M. Rietkerk and G. Sugihara, Early-warning  
492 signals for critical transitions. *Nature*, 461 (2009) 53-59
- 493 J. Sieber and J. M. T. Thompson, Nonlinear softening as a predictive pre-  
494 cursor to climate tipping. *Phil. Trans. Roy. Soc. A*, 370 (2012) 1166-1184
- 495 D. Sornette and A. Johansen, Large financial crashes. *Physica A*, 245 (1997),  
496 411-422.
- 497 P. Sura, M. Newman, C. Penland and P. Sardeshmukh, Multiplicative noise  
498 and non-Gaussianity: A paradigm for atmospheric regimes? *J. Atmos.*  
499 *Sci.*, 62 (2005), 1391-1409.
- 500 A. J. Veraart, E. J. Faassen, V. Dakos, E. H. van Nes, M. Lürling and  
501 M. Scheffer, Recovery rates reflect distance to a tipping point in a living  
502 system. *Nature*, 481 (2012), 357-359.
- 503 P. Wadhams, Arctic ice cover, ice thickness and tipping points. *J. Human*  
504 *Environment*, 41 (2012), 23-33.
- 505 C. Wissel, A universal law of the characteristic return time near thresholds.  
506 *Oecologica*, 65 (1984), 101-107.

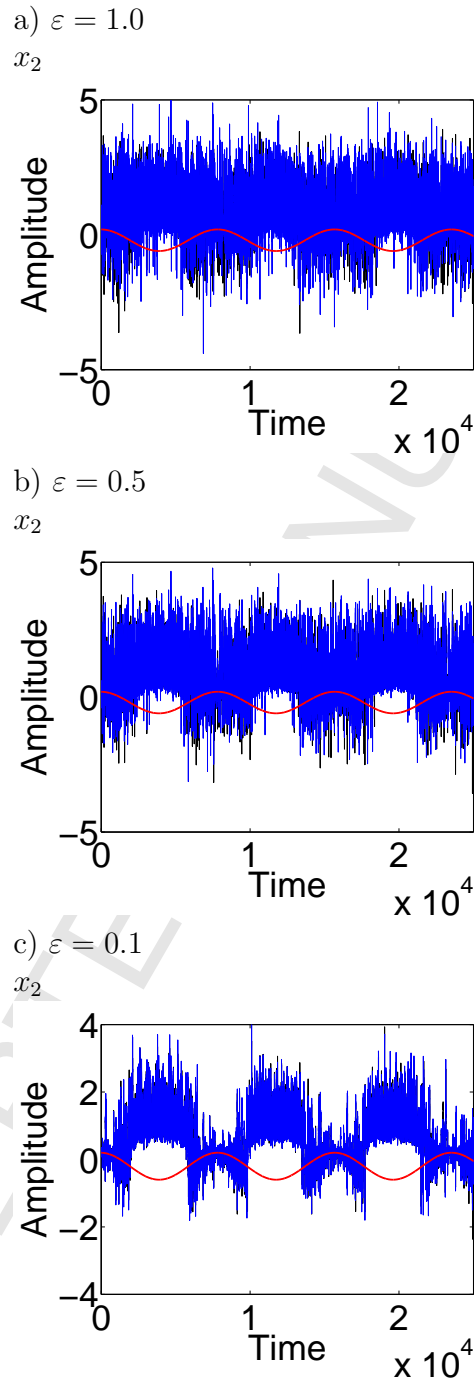


Figure 1: Section of one time series realisation for model simulations with forcing of resolved mode  $x_1$ : Black line: full dynamics; blue line: reduced dynamics; red line: Forcing.

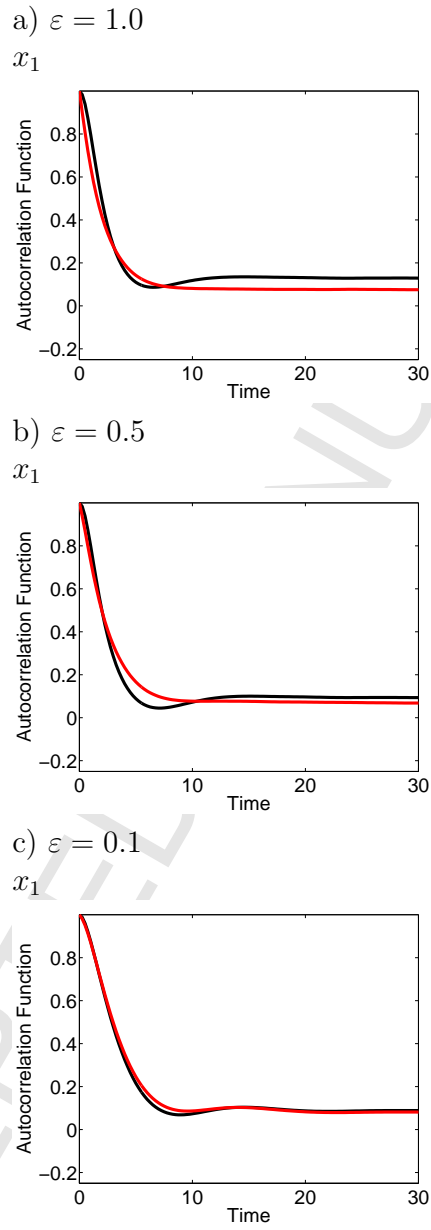


Figure 2: Autocorrelation function of model simulations with forcing of resolved mode  $x_1$ : Black line: full dynamics; Red line: reduced dynamics.

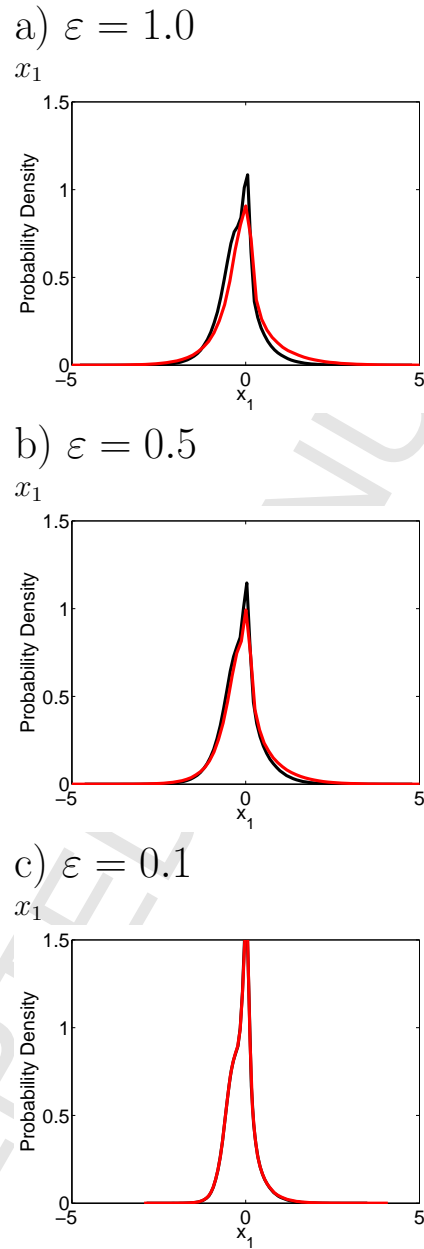


Figure 3: Marginal PDFs of model simulations with forcing of resolved mode  $x_1$ . Black line: full dynamics; Red line: reduced dynamics.



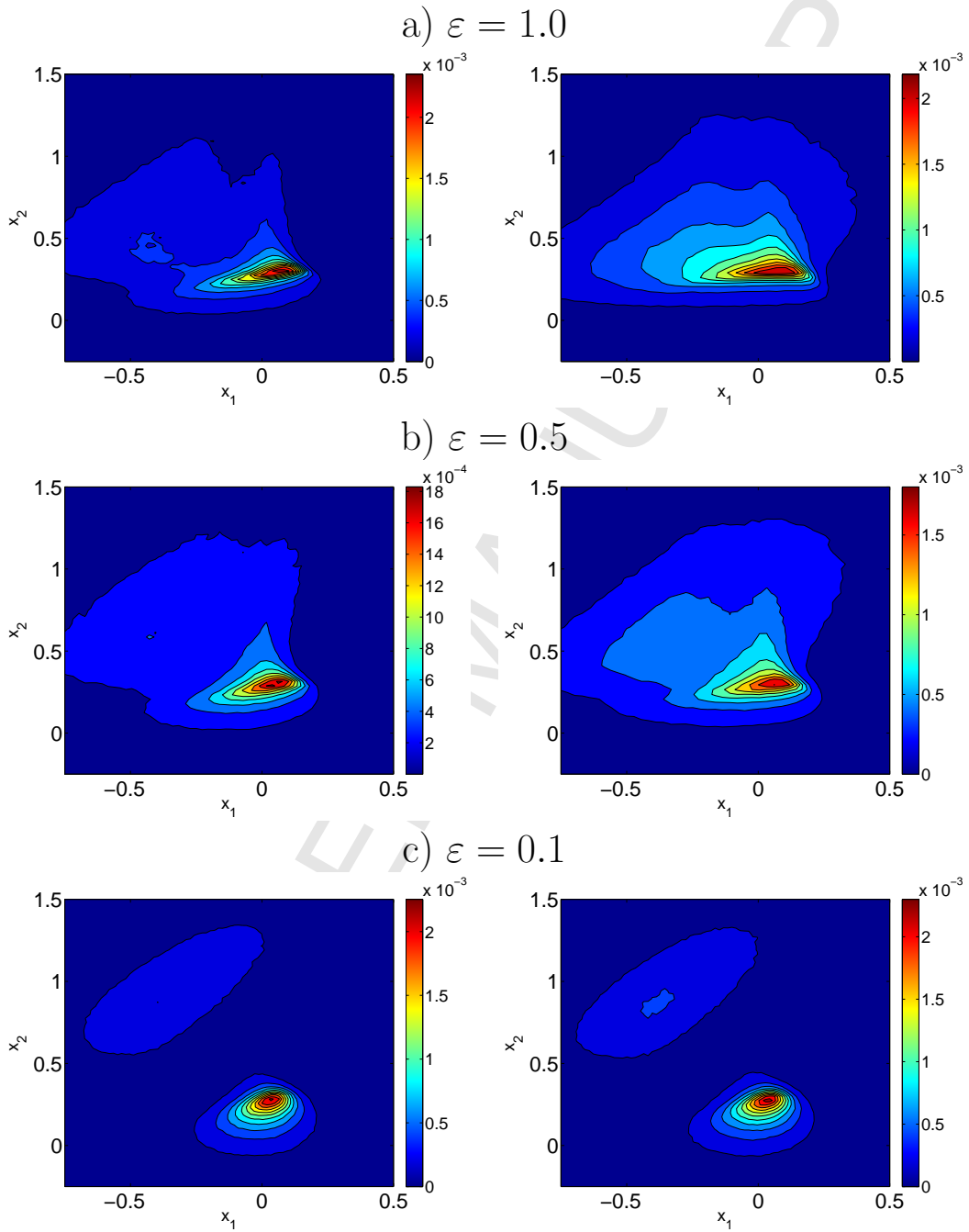


Figure 4: Joint PDFs of model simulations with forcing of resolved mode  $x_1$ : a)  $\varepsilon = 1.0$ , b)  $\varepsilon = 0.5$ , c)  $\varepsilon = 0.1$ . Left column: full dynamics, Right column: reduced dynamics.

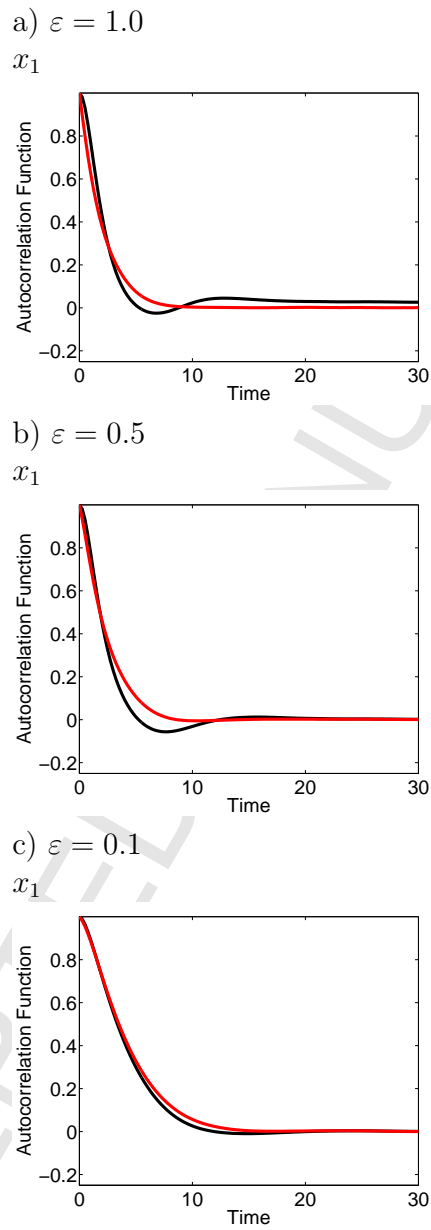


Figure 5: Autocorrelation function of model simulations with forcing of unresolved mode  $x_3$ ; Black line: full dynamics; Red line: reduced dynamics.

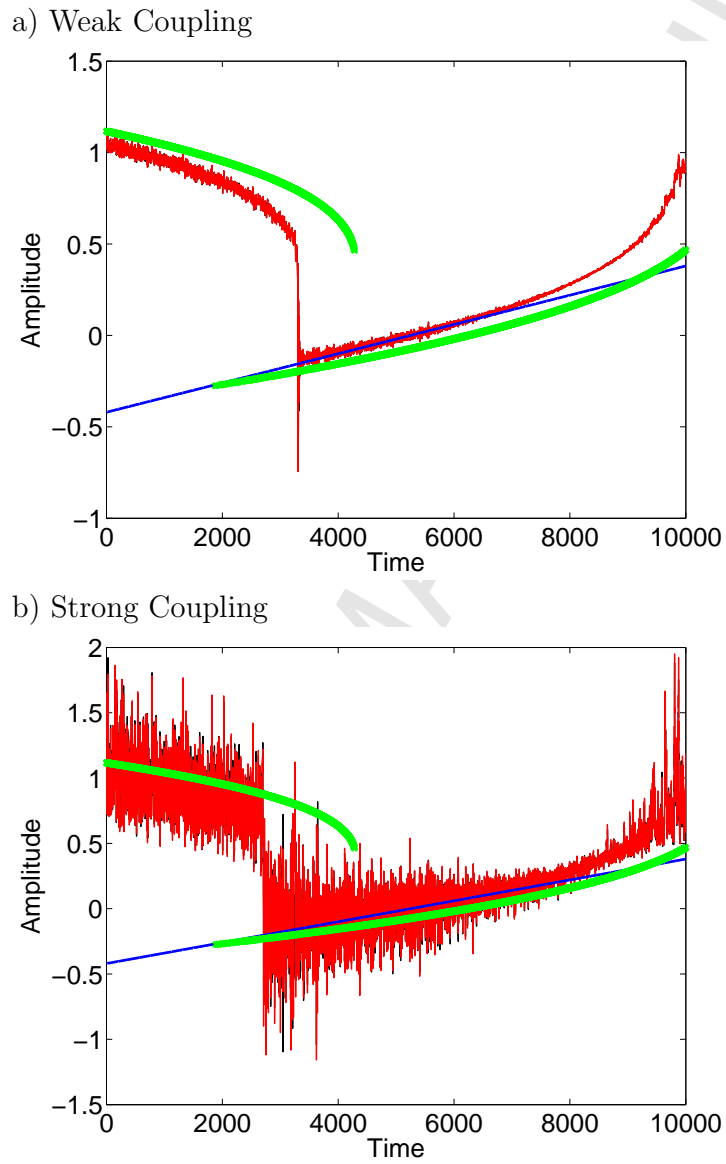


Figure 6: Time series of full dynamics (black line), reduced dynamics (red line) and forcing  $F_1(t)$  (blue line) for  $\varepsilon = 0.1$ . The green lines indicate the equilibrium solutions of the nonlinear deterministic system (17).

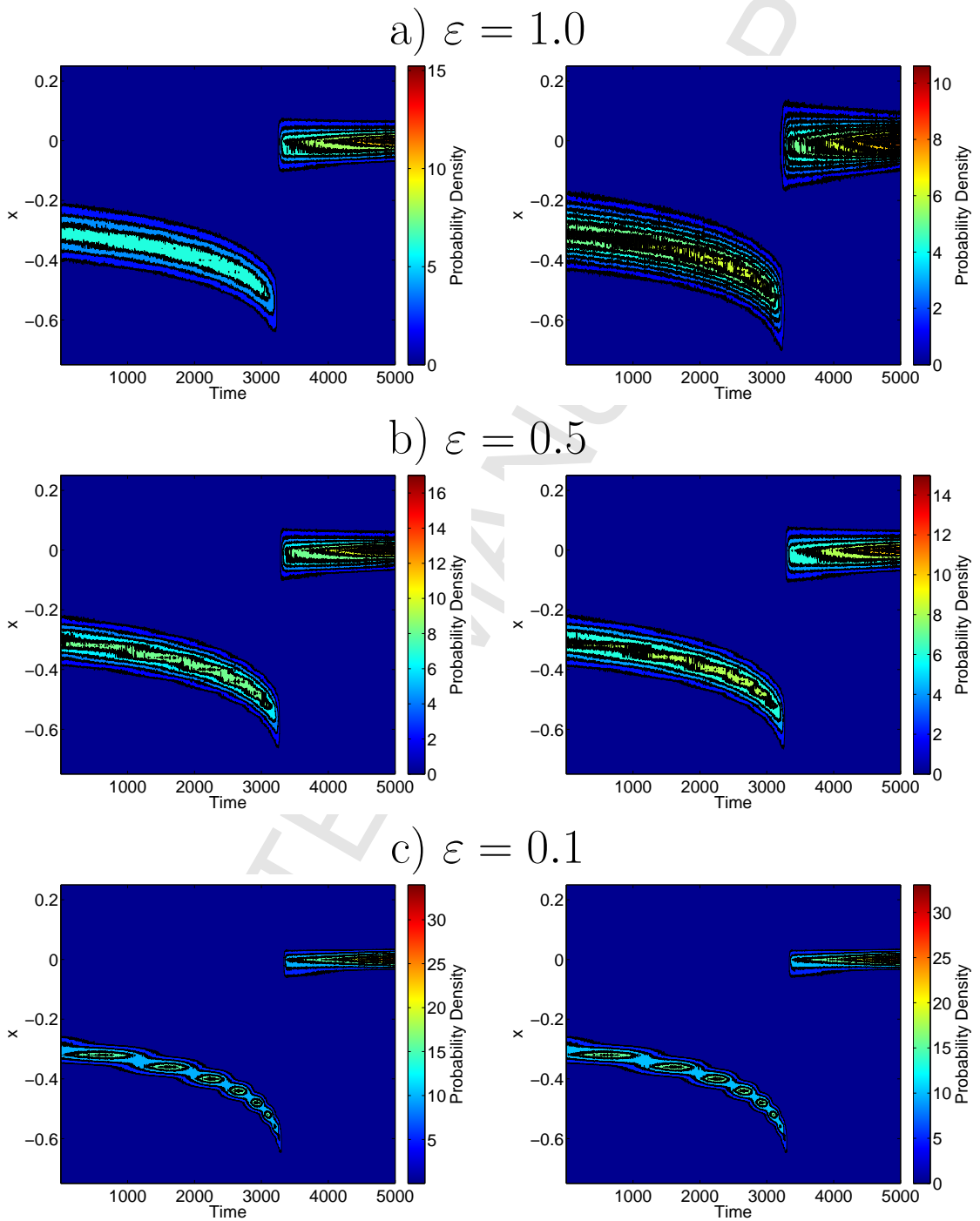


Figure 7: Marginal PDFs of  $x_2$  from 5000 member ensemble at time  $t$  for weak coupling of resolved and unresolved modes. Left column: full dynamics; Right column: reduced dynamics.

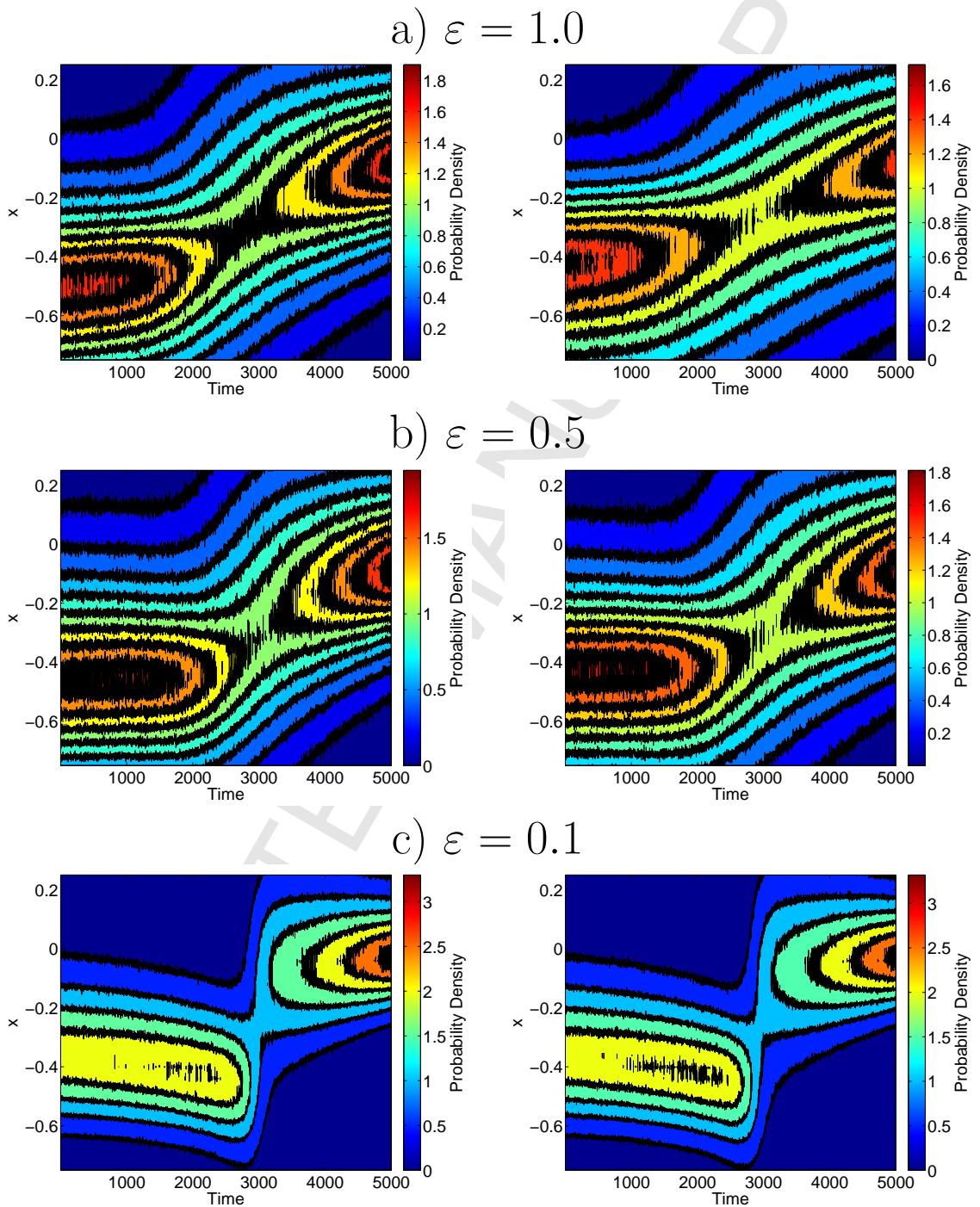


Figure 8: Marginal PDFs of  $x_2$  from 5000 member ensemble at time  $t$  for strong coupling of resolved and unresolved modes. Left column: full dynamics; Right column: reduced dynamics.

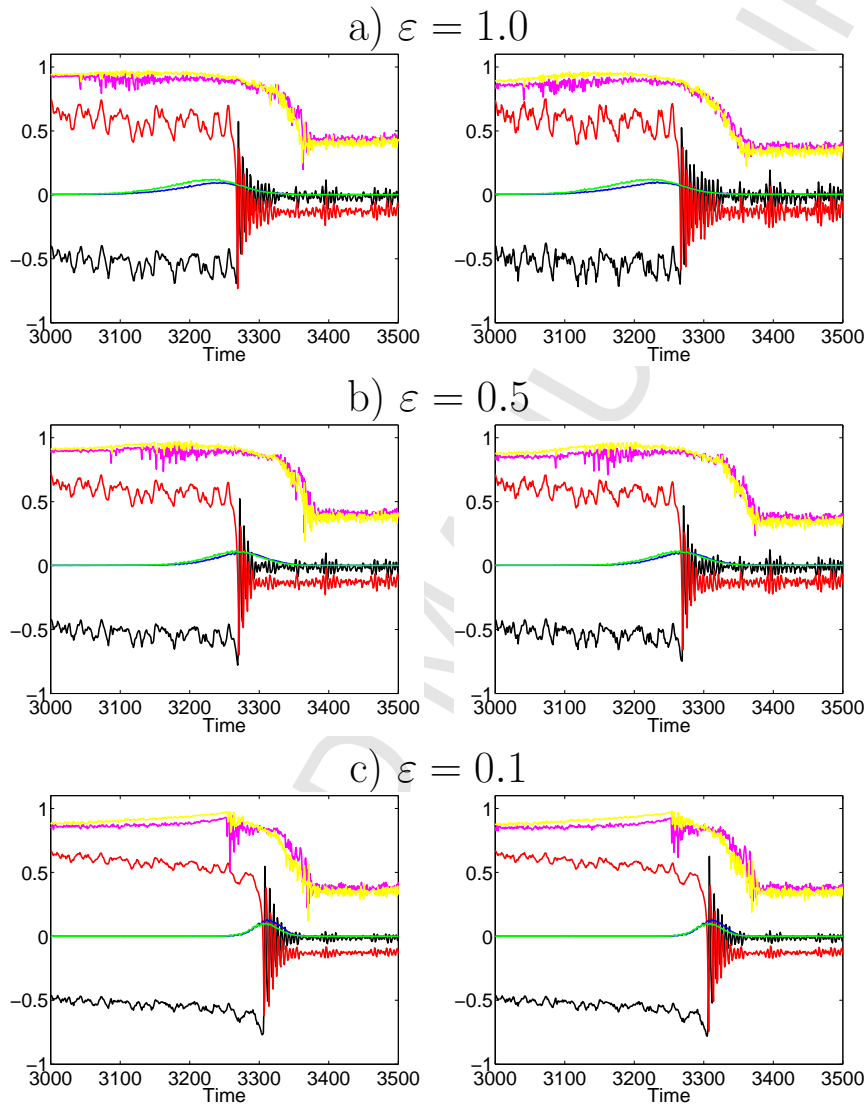


Figure 9: Ensemble indicator for weak coupling simulations. Left column: full dynamics; Right column: reduced dynamics. Black line:  $x_1$ , Red line:  $x_2$ , Blue line: ensemble variance of  $x_1$ , Green line: ensemble variance of  $x_1$ , Magenta line: ensemble AR(1) indicator of  $x_1$ , Yellow line: ensemble AR(1) indicator of  $x_2$ .

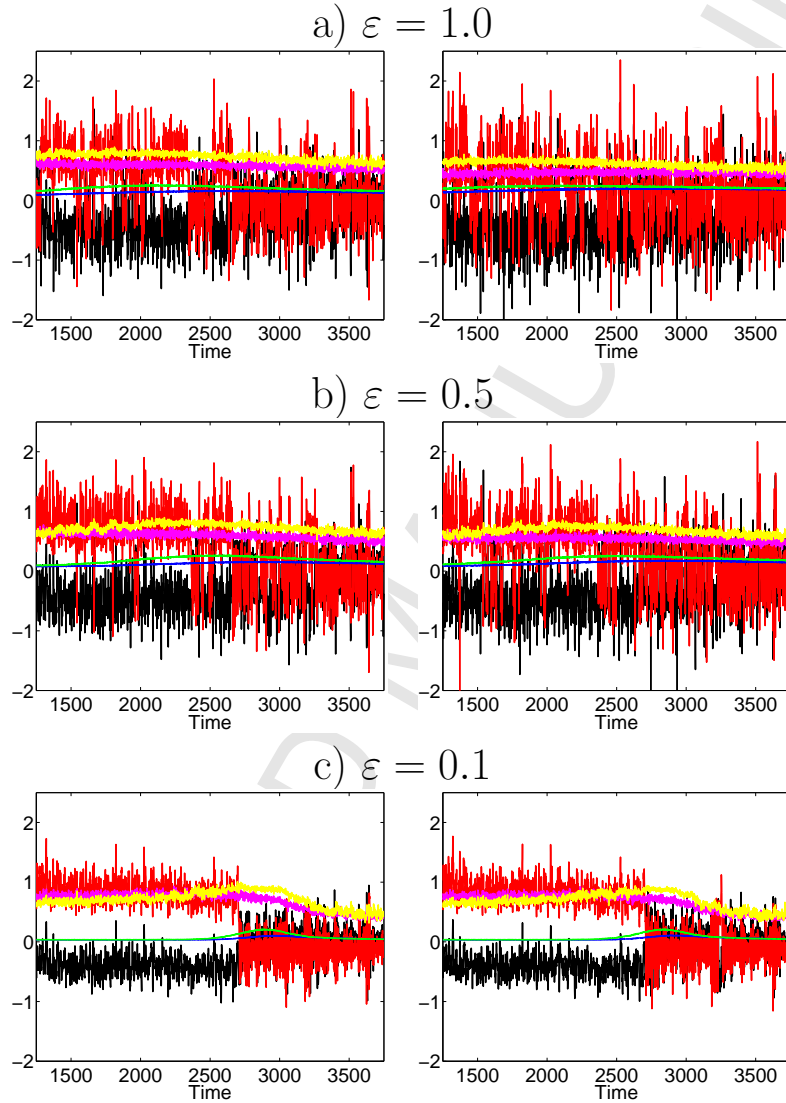


Figure 10: Ensemble indicator for strong coupling simulations. Left column: full dynamics; Right column: reduced dynamics. Black line:  $x_1$ , Red line:  $x_2$ , Blue line: ensemble variance of  $x_1$ , Green line: ensemble variance of  $x_2$ , Magenta line: ensemble AR(1) indicator of  $x_1$ , Yellow line: ensemble AR(1) indicator of  $x_2$ .

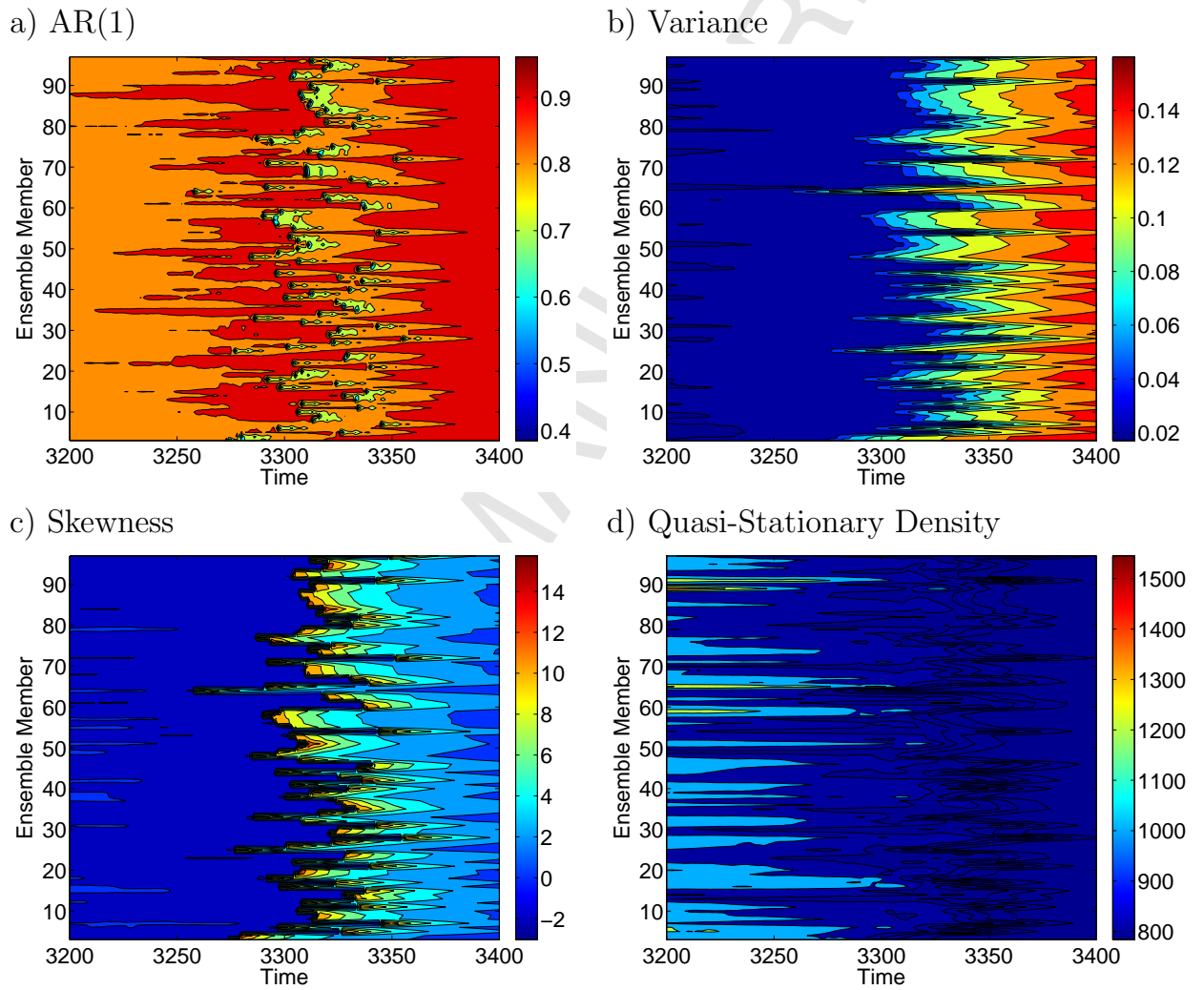


Figure 11: Tipping indicators applied to linearly detrended data from the full dynamics simulations over moving windows of length 1000 time units.