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Chemical Quality of Ground Water

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NEAM AND AND: **U.S. Composed CARTY** U.S. GEOLOGICAL SURVEY

WATER RESOURCES DIVISION HONOLULU, HAWAII

Report R48

Prepared by the

UNITED STATES GEOLOGICAL SURVEY

in cooperation with Division of Water and Land Development DEPARTMENT OF LAND AND NATURAL RESOURCES State of Hawaii

Chemical Quality of Ground Water in Hawaii

Report R48

By LINDSAY A. SWAIN

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UNITED STATES GEOLOGICAL SURVEY

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> *HONOLULU, HAWAII* May 1973

PROPFPTY OF U.S. GEOLOGICAL SURVEY .P.Q. & i412, HURON, SOUTH DAKOTA

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Introduction

As the quantity of water needed in certain localities of Hawaii is rapidly approaching the quantity of usable water available, identification and protection of the quality of existing and potential water supplies are becoming ever more critical. Certain factors are already identifiable as problems affecting the quality of ground water in Hawaii: (1) overdevelopment leads to salt-water encroachment by local upconing; (2) agricultural and industrial uses and discharges deteriorate both present and potential water supplies; and (3) the increasing practice of underground disposal of domestic wastes, with insufficient information to protect the water resource creates a potential hazard to developed and future supplies.

PURPOSE AND SCOPE

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This study was begun in mid-1971 by the U.S. Geological Survey in cooperation with the Division of Water and Land Development, Department of Land and Natural Resources, State of Hawaii. The purpose of the study was to identify present ground-water quality conditions, both natural and as affected by usage, to show readily identifiable changes in quality with time and usage, and to relate water-quality conditions to the development of ground-water supplies.

The study included a thorough research of all published data regarding chloride levels and complete chemical analyses for ground water in Hawaii; acquisition of data from local agencies collecting ground-water quality information; compilation of extensive unpublished water-quality data in U.S . Geological Survey files; coding of all data, where possible, for use in computer storage and retrieval systems; analysis of the collected and compiled data; and development of preliminary conclusions regarding the effects of pumpage and land use on water quality.

The report presents the results of a 2-year study of ground-water quality for the entire State of Hawaii.

PREVIOUS WORK

Concern for the quality of ground water in Hawaii dates back more than 85 years. The initial recorded sampling for chemical content of ground waters in Hawaii was made in early 1886 by Professor Lucius VanSlyke of Oahu College with the intent of determining water supplies of good, bad, or doubtful quality (Honolulu Water Commission, 1917).

Most water-quality studies of Hawaiian ground-water supplies are concerned primarily with chloride concentrations. Only a limited number have dealt with total chemical quality. Stearns and Vaksvik (1935) published the first extensive chemical analyses for Oahu. Bryson (1951) made an interpretive study of water quality in the Honolulu area. In 1961, Mink discussed the chemical characteristics of the intruded saline water underlying the fresh-water lens. Mink (1962) also discussed the effects of irrigation return water on the concentration of silica and nitrate in the ground water. Other studies on silica concentration in ground water and the effects of irrigation return water on ground-water chemistry were made by Davis (1969) and Tenorio and others (1969, 1970), respectively.

ACKNOWLEDGMENTS

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Special appreciation for cooperation and assistance during this study is extended to employees of the Honolulu Board of Water Supply, Departments of Water Supply of Maui, Hawaii, and Kauai Counties and the Division of Water and Land Development, State Department of Land and Natural Resources. Particular appreciation is expressed to the State of Hawaii Department of Health, especially Mr. William Wong, for his assistance in identifying much of the water-quality data. Numerous sugar and agricultural companies also graciously supplied chloride information from their files .

Factors Determining Ground-Water Quality

Ground-water quality is initially the result of effects or stresses caused by the various aspects of the hydrochemical cycle as the water moves through the system. Environmental factors creating the initial water quality are: (1) the chemistry of rain; (2) the effects of vegetation, soil, and geologic units on the water as it infiltrates; and (3) the results of mixing with saline water at depth within the aquifer. Further changes are caused by the activities of man, either by increased stress placed on the existing system or by simple addition of chemicals to the zone of infiltration.

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The source of almost all fresh water in Hawaii is rainfall; therefore, the chemical quality of rainwater is the genesis of ground-water chemistry in Hawaii.

Rainwater originates as water vapor from the ocean condensing upon salt nuclei, which are also of oceanic origin. The various ionic ratios in the rainwater are, therefore, quite similar to that of ocean water (Visher and Mink, 1964) . Slight differences in the ratios that exist probably result from additional chemicals of atmospheric, industrial, or even domestic origin.

The atmosphere may contribute gases, such as nitrogen in the form of nitrate and ammonium, carbon dioxide and oxygen. Industry contributes particulate matter from a wide spectrum of chemical constituents. Agricultural contributions may be windblown dust and particulate matter in the smoke from burning fields. Urban areas and domestic activities contribute numerous particulates to serve as nuclei.

The data-base for chemical quality of rainfall in Hawaii is quite deficient. The few existing analyses were made by Eriksson (1957, p. 520) for Project Shower on the island of Hawaii. Citing Eriksson's data and chloride-only analyses made by the U.S. Geological Survey, Visher and Mink (1964, p. 82) made additional interpretations of rain chemistry in a detailed study of southern Oahu. Analyses of rain chemistry are shown in table 1. Sea-water origin is apparent from the significantly greater concentrations of sodium and chloride, the major sea-water components.

The data ignore the existence of accumulated salts upon the land from interception of the daily aerosol or "dry rain" conditions in areas exposed to strong trade winds. The data available on water quality for high-level sources, however, indicate that the maximum natural-chloride concentration will usually not exceed 25 mg/1 (milligrams per liter) in Hawaii.

The relatively low pH of the rainfall indicates slightly acid conditions. This would cause the water to react with the rocks to bring about chemical weathering.

ROOT ZONE OF VEGETATION

The first terrestrial strata to have an effect upon the chemistry of the slightly acid infiltrating rainwater is the root zone of vegetation. As a result of plant respiration, carbon dioxide $(CO₂)$ is trapped in the pore space of the root zone. Upon contact with the infiltrating water, carbonic acid (H_2CO_3) is formed:

 $H₂O + CO₂ \rightleftharpoons H₂CO₃$

Passage through this zone thus adds carbonate (CO_3) and bicarbonate (HCO_3) ions to the infiltrating ground water.

GEOLOGY

The Hawaiian Islands are principally shield volcanoes of olivine basalt, which, in limited areas, may be capped by volcanics of either higher silica content or more ultrabasic character. Calcareous and alluvial deposits overlay these volcanics in some areas.

As chemical weathering plays a greater part in rock alteration than mechanical weathering in Hawaii, rainwater undergoes significant chemical changes depending upon the rock type and structure of the geologic unit through which it infiltrates.

Island Volcanics. The mineralogical composition of most island volcanics is such that the major minerals are readily decomposable. According to · Macdonald and Abbott (1970, p. 153), the principal minerals that produce soluble products are olivine, anorthite, nepheline, pigeonite, augite, and albite. From decomposition of these minerals, the products released to solution are: MgC03, FeC03, NazC03, and SiOz. The abundance of each will depend upon the chemical composition or ease of decomposition of each mineral. Thus, from the volcanics the infiltrating water acquires calcium, magnesium, sodium, silica, and iron, although the amount is limited.

Table 1. Average composition of rainfall caught at high- and low-rainfall localities on Hawaii and Oahu (Constituents in parts per million)

Reprinted from Visher and Mink, 1964, p. 82.

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Calcareous Deposits. Although information about water quality within the coralline reef or other calcareous deposits is limited, Visher and Mink (1964, p. 116) have noted the major soluble products from the decomposition of these rocks. These products include higher proportions of carbon dioxide (CO2), bicarbonate (HCO3) and calcium (Ca) from the calcareous rock than from other rock types. The proportions are also higher than in ocean water. In addition, slightly increased concentrations of magnesium may result from minor impurities in calcareous rock.

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Because of the significantly increased bicarbonate concentrations in the calcareous rock zone. analyses with bicarbonate ion concentrations greater than 140 mg/1 (that of ocean water) may indicate solution of calcium carbonate rock. Exceptions to this condition may exist when there is considerable recycling of irrigation water or possible volcanic gases being released into the ground-water source.

Alluvial Deposits. Because alluvial deposits are limited in extent throughout the islands and the unit is generally of low permeability, this rock unit is a very minor source of fresh water. In addition, where drilled wells pass through the unit to reach the volcanic aquifer beneath, the entire thickness of alluvial rock is usually cased off. Thus, the information on quality within this aquifer is extremely limited. However, the chloride-only information that is available shows a great variation of values which is apparently due to permeability differences affecting the movement of sea water into the rock unit. Visher and Mink (1964) indicate that calcium and magnesium concentrations are increased by cation exchange within the sediments.

SEA WATER

The aquifers of Hawaii exist in an island environment and are everywhere surrounded by sea water. Unless these highly permeable waterbearing rocks are capped by low-permeability material or contain low-permeability dikes, the ocean may freely penetrate wherever the aquifer is exposed below sea level and fluid pressure gradients are such as to permit it. Where fresh-water heads are at or near sea level, tidal fluctuation and other head variations tend to create a zone of mixing and broad zone of transition from ocean water to fresh water.

Ghyben-Herzberg Lens. According to the Ghyben-Herzberg principle, density differences between fresh and salt water in a static system will allow infiltrating fresh water to float upon salt water. The ratio of densities is such that for every unit of head of fresh water above sea level, the saltfresh water interface will be depressed approximately 40 times that unit below sea level.

However, as Hubbard (1940) pointed out, the Ghyben-Herzberg lens is a dynamic, not static. system. Thus, as movement occurs within the water bodies, the fresh and salt waters mix and change the interface from a sharp line to a gradational zone of transition. The thickness of the zone depends upon the magnitude of mixing caused by tidal influences, upconing of the salt water from pumpage, and variation in recharge.

Chloride as a Sea-Water Indicator. Because the chloride concentration is quite high in sea water, chloride has long been used as the principal indicator of sea-water intrusion. It takes little more than 1 percent sea water to reach the 250-mg/1 chloride limit for drinking water. Thus, the determination of the chloride concentration of ground waters aids in determining the probable extent of sea-water intrusion.

The utility of chloride as a sensitive indicator of sea-water intrusion is enhanced by its inherent qualities. It will not readily enter into chemical reaction with other chemical constituents in the ground-water environment, nor is it subject to appreciable ionic exchange. Also, there are no known major sources of chloride within the lavas of Hawaii.

Chemical Characteristics of Intruded Water. Each aquifer unit has fresh water with its own inherent quality. However, when the influence from salt-water intrusion becomes prominent, the high chemical concentrations of the sea water overshadow the fresh-water concentration. In some instances, the intruding sea water will also undergo chemical changes as it moves into the aquifer from the ocean.

Mink (1961) noted that when sea water intrudes a volcanic aquifer, which is overlain by an alluvial caprock, the intruded water has a chemical character significantly different from that of ocean water. His analysis of water from a deep test hole in the Pearl Harbor area of Oahu showed more calcium and magnesium and less sodium and potassium than would exist in ocean water diluted to the chloride concentration of the deep-well sample. This change has been attributed to a cation exchange as the ocean water penetrates the ocean-bottom sediments and areas of permeable caprock, exchanging the high concentrations of sodium for calcium and magnesium in the clay structure of these deposits.

In addition, the sulfate and bicarbonate ions are also less than would be expected from simple dilution of sea water. The lesser sulfate is believed to result from sulfide reduction in the bottom sediments (Mink, 1961). The anomalous decrease in bicarbonate is hypothesized to occur as a result of pressure loss as water under aquifer pressure reaches the reduced pressure of the well, with the resultant release of carbon dioxide and deposition of calcium carbonate.

INTRODUCED CONTAMINANTS

The numerous and varying activities of man have, in some instances, altered the chemical character of the ground-water bodies beyond that acquired from environmental sources. Man's effects fall into three broad categories: agriculture, industry, or domestic waste disposal.

Agriculture. Studies on Maul and Oahu have described the water-quality change effects of irrigation in sugarcane areas (Tenorio and others, 1970; Takasaki and Valenciano, 1969). These changes have been attributed to both stress on the fresh-water lens by pumping, and to the addition of chemicals as a by-product of fertilizer and other surface chemical application.

Intensive pumping of the basal fresh water has been shown to cause local upconing of the saline-water body, resulting in higher concentrations of most ions, especially chloride. Irrigation with this water of higher-chloride concentration will further increase the chloride concentration in the fresh-water lens, as excess irrigation water infiltrates back to the freshwater lens.

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Fertilizer and other surface chemical applications add soluble products to the ground-water bodies but, at present, none of these additive chemicals have been detected in excess of the limitations for drinking water. Nitrate and sulfate have been the most extensively studied of these additive products. These chemicals infiltrate from the irrigated fields, and often a concentration, which decreases with depth, is found accumulated above intruded saline water.

Nitrate (NO_3) has an average concentration in rainfall, in native fresh water and even in sea water of less than 1 .1 milligrams per liter (Mink, 1961). Where concentrations greater than this occur beneath areas of irrigation, the parameter may be considered a reliable indicator of return irrigation water.

Because sulfate is also a soluble product of fertilizer, it too is a good indicator of return irrigation water. However, sulfate concentration in sea water is about 2,600 mg/1, and high concentrations may simply indicate sea-water intrusion unless other constituents such as chloride are not in proper proportion.

Visher and Mink (1964) recognized that even though phosphate and potassium are highly concentrated in fertilizers, they are poor indicators of return irrigation water. Phosphate is immediately fixed in the soil as an insoluble product, and potassium is subject to ion-exchange reaction; the unconsumed amount usually remaining in the soil.

Tenorio and others (1970) noted that increased application and recycling of water in certain geologic environments w111 accelerate solution of specific minerals by the infiltrating water. This situation is prevalent in central Maui where high concentrations of bicarbonate are encountered in ground waters below heavily irrigated areas. Because the zone of infiltration in this area is composed largely of calcareous deposits, it is not known for certain whether the high bicarbonate values are the result of continuous recycling through the root zone of vegetation or whether from solution of the carbonate rock, or both.

Industry. Although various wastes, such as hydrocarbons, metals, detergents, and other chemical products are created by industrial activities, only very limited information is availabie regarding the influence of such industrial waste upon the ground-water quality in Hawaii.

In one recent study, Zaidi (1973) found the levels of organic carbon in Oahu ground water to be well below suggested limits for potable water, and that in the artesian aquifer beneath the industrialized Honolulu-Pearl Harbor area, the levels are quite low. This is most likely attributable to a thick, nearly impermeable alluvial caprock and the upward pressure in the artesian aquifer.

A recent reconnaissance of domestic groundwater supplies by the Honolulu Board of Water Supply found the mercury concentration of all the numerous ground-water sources sampled to be well below the recommended limit of 5 micrograms per liter.

No other extensive studies are known to have been made on other trace metals or possible industrial contaminants in Hawaii. However, as industry expands with population to areas beyond the caprock area and over the recharge areas for the islands' main aquifers, monitoring of the groundwater system for industrial contaminants will become important.

Domestic Waste Disposal. Solid and liquid domestic waste-disposal practices and their chemical effect on specific ground-water bodies have not been studied in great detail for Hawaii. Takasaki (1973) has inventoried the extent and type of domestic waste-disposal practices in the State, has identified present practices, and has delineated, generally, areas where the possibility for contamination may exist.

Other than significant nitrate concentrations in some cesspool areas, relatively little data exist to identify specific chemical effects from domestic waste-disposal practices in Hawaii.

The Ground-Water Quality Data Base

Since 1886, when the first chloride determinations were made on Oahu, various private and governmental agencies have been collecting chemical-quality information about ground-water sources in Hawaii. Since then many agencies and private organizations or individuals have collected chemical-quality data at various locations for a variety of purposes . Principal among those concerned were the agricultural interests who were the principal users of ground water for irrigation. As more ground water began to be used for domestic and municipal purposes, the Territorial, State, and City and County agencies became more involved. Chloride was the primary indicator of usability, and the great bulk of data contains only chloride determinations. Some uses, however, required more detailed knowledge of dissolved constituents, and a small body of data from complete analyses was developed.

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TYPE AND AMOUNT OF DATA AVAILABLE

As noted, by far the majority of all waterquality data consists of only chloride-concentration analyses . Although some sources may have only one chloride analysis available., others have more than 60 years of regular analysis for chloride. The amount of this data from the more than 1, 700 groundwater sources on Oahu alone is voluminous. Although less extensive, a great amount of data is also available for the other major islands. The chloride-only data base used herein was developed from analyses compiled during this study from various sugar and agricultural companies, county water departments, and the files of the U.S. Geological Survey.

In addition to the chloride-only data, more than 2, 000 complete chemical anslyses of groundwater sources within the State were obtained during this study. These analyses were made in the laboratories of the City and County water departments, State of Hawaii Department of Health, University of Hawaii, the U.S. Geological Survey, and a few private agencies. Many recent analyses with wide geographic distribution were available for the study through a concurrent ground-water data network program being carried on by the Geological Survey wherein a large number of sources were sampled and complete analyses were made.

EVALUATION OF THE DATA

Because the complete chemical analyses were made by different agencies, they vary greatly in precision owing to differences in sampling purpose and analytical technique. The State Department of Health and the Honolulu Board of Water Supply take . samples of domestic supplies to ensure that no chemical parameter exceeds recommended limits as set forth by the U.S. Public Health Service (1962) . The University of Hawaii made analyses during 1968, 1969, and 1970 on some irrigation sources to find a means of identifying irrigation return water in the ground-water bpdies. The U.S . Geological Survey has taken extensive samples on a state-wide basis, in order to evaluate the chemical quality of all ground-water resources in Hawaii.

For the collected data to be useful in evaluation of the total chemical-quality conditions, or to show changes for certain parameters, the data from all agencies had to be made comparable in content and precision.

The major content discrepancies existed because the methods of reporting constituents were not uniform. Some agencies reported the individual ion not as ion concentration, but as ion complexes or alkalinity values. In addition, the computed values of calculated solids, carbonate and noncarbonate hardness, and alkalinity were either not reported, or were determined by varying methods. Thus, to make these analyses comparable it was necessary to recompute them to their individual ion concentrations . Finally, the calculated values were recomputed, using nationally accepted methods as set forth by Rainwater and Thatcher (1960).

Each individual analysis was intensively examined to determine which analyses were of reliable precision. The initial step in identifyinq unreliable analyses was a check of the cationanion balance. To be truly representative of the actual chemistry of the sample, total cations in milliequivalents per liter (meq/1) should approximately equal total anions in meq/1. However, because analytical technique varies considerablY among analyzing laboratories, the balance is rarely exact. An arbitrary limit of cation-anion difference of 5 percent of the total was set as the level of analy sis precision acceptable for this study. To further complicate data evaluation, analyses made prior to

1940 sometimes reported sodium concentration as calculated from the cation-anion difference, rather than as determined directly. For these analyses checks were made by comparing the individual parameters in each analysis with values on other dates for the same source or with other sources in the same aquifer located nearby. By this check, analyses which contained gross variations from the normal were eliminated.

Following the checks on content and precision, approximately 900 analyses that met the qualifications for the study were processed. The data .are shown graphically and by numerical value on the accompanying plates, and are tabulated in Appendix A at the end of this report.

WELL-NUMBERING SYSTEM

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January 2001–2010
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The ground-water data-point identification system for Hawaii has recently been standardized, and is the system in use by the U.S. Geological Survey and the State of Hawaii, Department of Land and Natural Resources, Division of Water and Land Development. This system makes areal location of any ground-water data source easily obtainable from its identification number. Each island is divided into minutes of latitude and longitude, and each one-minute-square block is identified by a four-digit number; the first two numbers being minutes of latitude and the last two numbers being minutes of longitude. Within each one-minutesquare area, ground-water sources are numbered serially by a two-digit number, preceded by a dash following the four-digit block number. Thus, the

first source numbered in the one-minute-square block at 34 minutes of latitude and 26 minutes of longitude would be identified as 3426-01. Wherever possible, the serial number assigned is based on the chronological order in which the source was identified or (if a well) constructed. In a few instances where the water-quality analysis is from a combination of wells, the first well identified in the active battery will bear the serial identifier and this will be followed by the letter Z to identify a combination sample.

WATER-QUALITY TYPES

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In this report, the names used for waterquality types identify ground water of a particular chemical character and are applied as in the following examples: (1) "sodium chloride" designates a water in which sodium amounts to 50 percent or more of the cations, and chloride amounts to 50 percent or more of the anions, in milliequivalents; (2) "calcium magnesium chloride" designates a water in which calcium and magnesium are first and second, respectively, in order of abundance among the cations but neither individually amounts to 50 percent of the cations; (3) "calcium bicarbonate chloride" designates a water in which bicarbonate and chloride are first and second in order of abundance among the anions, as above; (4) "calcium magnesium sodium chloride" designates a water in which calcium, magnesium, and sodium are each of approximately equal percentage of the total cations .

Ground-Water Quality by Island

Although the magnitude of each chemical parameter may vary considerably, the basic conditions determining ground-water quality can generally be applied for each of the islands. From the large amount of data acquired during this study, the water-quality types, their distribution, and . even some conclusions as to specific causes of greatly varying conditions can be described for ground-water sources on each of the five major islands.

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Kauai, oldest and northernmost of the islands, is. the most geologically complicated. The island was formed from a major shield volcano of principally basaltic composition. The volcano later collapsed to form the State's largest caldera, and lava flows subsequently erupted on its periphery. Following a long period of quiescence and erosion, more recent and extensive lava flows occurred within the eastern two-thirds of the island (Macdonald and others, 1960) .

Ground-Water Zones. Kauai's ground water occurs mainly in the Koloa Volcanic Series and in the Napali formation of the Waimea Volcanic Series. Water in these volcanics may occur as perched water, dike-impounded water, or as a basal lens .

The perched water occurs mostly in the poorly permeable to moderately permeable Koloa Volcanic Series at lower elevations, where it is perched by ash or tuffaceous soil beds (Macdonald and others, 1960) .

The dike-impounded water is found mostly in the highly permeable Napali formation along the Napali Coast and in the Makaleha Mountains.

Basal water occurs in both rock units but owing to the complicated geology and widepermeability ranges of the lava flows, conditions are not favorable for large, well-developed, freshwater lenses .

Water-Quality Types by Zone. As depicted on plate 1, the overall chemical quality of ground water from 25 sources on Kauai is excellent. However, previous hydrologic studies have noted that chloride concentrations exceeding 250 mg/1 exist for Kauai, primarily along the western coastal plain. An examination of the available analyses, coupled with available hydrologic information on the water sources, shows a vague pattern of water-quality

variation according to area and aquifer type.

The two analyses of perched water from wells 1225-01 and 5823-01 are calcium bicarbonate and magnesium sodium .bicarbonate types, respectively. Each analysis shows a low dissolvedsolids concentration . The analyses of dike-zone water from wells

1331-01, 0623-01, and 0623-02 show that all three are of the magnesium sodium bicarbonate type with low dissolved-solids concentrations.

The remaining analyses, which are all believed to be from basal-water sources, vary considerably in water-quality type. In addition, they each show higher dissolved-solids concentrations than those for dike-impounded or perched waters. The major anion in all but five of these samples is bicarbonate. Where sea-water intrusion is evident, as in well 0120-01, chloride is the dominant anion. The increased chloride of this well is accompanied, however, by a significant percentage increase of calcium and magnesium and a decrease of percentage of sodium over what would result from simple mixing. Thus, it is interpreted that the intruding sea water has been altered by cation exchange.

The significant differences in water quality in wells close together, such as 5534-02 and 5534-03, demonstrate the great variabilities encountered in the water bodies of Kauai. In this example, the variability may be caused by an effective groundwater barrier of Hanapepe Stream Valley fill between them. In other localities, differences may result because of differing rock types, because of local geologic structural feature, or from varying landuse practices in the recharge areas.

OAHU

Oahu is the third largest of the islands. It is the population and industrial center of the State with the capital city, Honolulu, on its southern shore. The island had its origin from two volcanoes, which now exist as the eroded Koolau and Waianae Mountain Ranges interconnected by a central plateau. Oahu's very irregular coastline exhibits the effects of Pleistocene relative sea-level fluctuations, with coralline and beach deposits existing far inland from the present-day shoreline. In addition, Oahu has by far the most extensive alluvial deposits of any of the islands. Thus, the ground water of Oahu exists in three major aquifers of great extent: the volcanic bedrock, calcareous sedimentary deposits, and noncalcareous sedimentary material.

Water in the Noncalcareous ·Sedimentary Materials. Noncalcareous sedimentary material in Oahu is shown on plate 2, and consists mostly of talus and alluvium derived from the volcanic bedrock or the reworking of older talus and alluvium. Stearns and Vaksvik (1935) divided this material into two units, the Pleistocene consolidated sediments (older alluvium) and the Holocene unconsolidated sediments (younger alluvium) . The older alluvium, which is moderately to well consolidated, is poorly permeable. The younger alluvium is more permeable than the older, but is usually less than 20 feet (6.1 meters) thick and situated above the water table. At and near the shore it is usually saturated with brackish to salty water. ·On plate 2 the two alluvial units are combined and their distribution is shown as the combined unit.

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Only a small number of wells on Oahu tap the alluvial unit for water. Most of these exist only because of failure to develop a suitable supply in deeper, more permeable lava flows or, in some coastal areas, in coral limestone. In 1962, about 25 shallow test borings along the west coast of Oahu, and about 10 test borings in north-central Oahu, were drilled to investigate ground water in the sedimentary sectons along the coasts. Much of the information on plate 2 was taken from findings of that investigation.

The relatively low permeability of the noncalcareous sedimentary materials helps to retard sea-water intrusion. Therefore, this aquifer generally contains fresher ground water than that found in the more permeable lava flows and coralline limestone along coastal areas. The range of chloride contents for sources within the alluvial deposits during 1958-70 is shown on plate 2. Even though individual well yields are generally low, the aquifer supplies much needed water for small truck farms in the Waianae area.

Only a limited number of chemical analyses are available for water in the alluvial deposits. However, Visher and Mink (1964, p. 112) have noted that cation exchange within the sediments is significant, as shown by magnesium and calcium concentrations, which are increased from that of water in other aquifers .

Water in the Calcareous Sedimentary Materials. Calcareous sedimentary material in Oahu is shown on plate 3, and consists mostly of coral limestone with minor amounts of dune and beach sand. Logs of dr1lled wells indicate coral limestone to a depth of at least 1, 178 feet (359 meters) below sea level in at least one locality. . Most of the limestone, however, is found in the interval from about 30 feet (9 .1 meters) above sea level to about 250 feet (76 meters) below sea level.

According to Stearns and Vaksvik (1935, p. 169) , lithified dune sand from ancient beaches has a maximum thickness of about 125 feet (38

meters) in Oahu. Most of this material lies .above the water table, as does unconsolidated dune sand blown inland from existing beaches. Unconsolidated beach sand rarely exceeds 25 feet (7. 6 meters) in thickness and usually contain brackish or highly saline water. The exposures of limestone, dune sand, and beach sand are combined and shown as a single unit on plate 3.

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The most important aquifers of this unit are in the coral limestone. Ground water pumped from the unit is the major source of water for sugarcane irrigation in the Ewa area west of Pearl Harbor. Another important use of the coral limestone aquifers is for cooling water for air-conditioning units in Honolulu. The calcareous unit is also an important source of water for irrigation of small truck farms in the Waianae area.

The high permeability of most of the calcareous deposits permits extensive sea-water intrusion which, in turn, greatly limits the usage of water from this unit. The chloride levels for sources 'within this unit are shown on plate 3.

During recent years, the coral limestone and unconsolidated sands have been used increasingly as receptacles for liquid waste. Initially, the materials were discovered to be good, reliable, recipients of cesspool seepage. Later, liquid waste of this and other kinds was injected *by* means of trenches, dug wells, and drilled wells. The liquid waste included air-conditioning water return, car-wash effluent, flood water, and drainage water. The practice of disposing of effluent from private sewage-disposal systems into limestone and beach deposits has been increasing, owing to increasing development of residential subdivisions and hotels remote from existing sewers .

Only three complete water-quality analyses from calcareous sedimentary aquifer were available for this study. The analyses, as shown on plate 4 (labeled "limestone"), show the water to be a sodium ·chloride type, in both the Waialua and Waimanalo wells. However, water from the limestone aquifer in Pearl Harbor is a sodium magnesium chloride type.

One distinctive characteristic common to all of these analyses is the high bicarbonate. Values range from 203 to 332 mg/1, which greatly exceeds that of sea water and indicates dissolution of the CaCO₃ rock.

Water in the Volcanic Bedrock. The term "volcanic bedrock" for Oahu, as used in this report, applies only to rocks associated with the initial phase of mountain building, caldera collapse, and caldera filling. The rocks included are those of the Waianae, Kailua, and Koolau Volcanic Series, as described by Stearns and Vaksvik {1935) . Rocks of the Honolulu Volcanic Series of the late Pleistocene and Holocene time constitute only a fraction of the total rock volume and are not included.

Most of Oahu's ground water occurs in the

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Compiled by L.A. Swain, 1973

volcanic bedrock as dike-impounded water in rift zones and as basal water outside the rift zones . Lava flows tend to be more thickly bedded and are dike-intruded within the rift zones, but outside the zones, lava flows are generally thinbedded and dike-free.

Water-quality types of high-level water areas: . As shown on plate 4, most of the water sources in Oahu's three high-level water areas (Koolau Range, Waianae Range, and the central plateau high-level water area) have quite similar chemical characteristics, Sea water is effectively barred from these areas by high fresh-water "heads, by the dikes as impermeable boundaries, or by other impounding geologic structures . It appears that the chemical character of these waters is derived almost entirely from rainfall interacting only with the rocks within the areas. The water-quality type for all three areas is predominantly sodium magnesium bicarbonate, with very low concentrations of chloride and dissolved solids .

A few exceptions to these types exist on the northwestern edge of the Waianae Range, and in the northern coast of the Koolau Range. In these two areas, the water is affected by sea-water intrusion. The chemical character of the water in these areas is more like that of basal water than dike water.

One important area, with a water quality distinctively different from that of dike water. exists within the central portion of the Waianae Range. At two locations, wells 2508-02 and 2808-01, the water has a significantly greater dissolved-solid concentration than that of surrounding sources . The water for well 2508-02 is a sodium sultate type, and that for well 2808-01 is a magnesium chloride type. These waters are markedly different from that found on any of the other islands . The anomalies may be related to a geothermal hot spot in this area. Water temperatures of 26.7° C for well 2808-01 at an elevation of 440 feet (134 meters) and 28. 4°C for well 2508-02 at a lower elevation would tend to support this possibility.

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Water-quality types of north-central basalwater area: The north-central area extends from the northeastern edge of the Waianae Range to the western edge of the Koolau Range north of the central plateau high-level water area. The basal water in this area can be subdivided into three subzones based on water levels and ground-water boundaries (Rosenau and others, 1971).

The westernmost third of the area has artesian water levels from 10 to 24 feet (3,0-7.3 meters) above mean sea level (msl) , with a relatively thick caprock. The water in this area is either a sodium magnesium chloride type or a magnesium sodium chloride type, The dissolved-solids concentration ranges from 280-590 mg/1.

The middle subzone, with less caprock, has artesian water levels averaging about 11 to 13 feet (3. 4-4.0 meters) above msl. The water in this area is a sodium chloride type. The sodium plus chloride concentrations in water in this subzone is between 50 and 75 percent of the total concentration of the major ions,

The subzone northeast of the valleyfill boundary has water-table conditions with water levels less than 5 feet (1, 5 meters) above msl, and relatively little caprock. The two analyses for water in this subzone, shown on plate 4, have greater than 900 mg/1 dissolved-solids concentrations. The water is also a sodium chloride type but the sodium plus chloride concentration is greater than 75 percent of the total ion concentration.

Water-quality types of the Kahuku basalwater area: The Kahuku basal-water area is located along the northeastern edge of the Koolau Range. Although there is no absolute ground-water boundary to separate them, two distinct water-quality types are evident.

The southernmost half of the area has artesian water levels from 18 to 22 feet {5. 5-6.7 meters) above msl because of an alluvial caprock. The water in this area is primarily a sodium magnesium chloride type. The dissolved-solids concentration of these waters is less than $235 \text{ mg}/1$.

The northern half of the area has extensive calcareous deposits both in the upland recharge areas and intermixed with alluvial caprock. Ground water in this area has considerably greater dissolvedsolids concentration than that in the southern section. The water here is mainly a magnesium calcium chloride type. There appears to be extensive cation exchange present. The high pumpage rates and varied geology of this area definitely affect the water quality.

Water-quality types of the southern Oahu basal-water area: The numerous analyses from greatly varying depths and. hydrologic conditions make it difficult to generalize as to water-quality types in the southern Oahu area.

However, the dissolved-solids concentrations of water samples from the Honolulu area are less than in the Pearl Harbor area (insets, plates 4A and 4B). In addition, the ground water of the Honolulu area is primarily a sodium chloride type, while those waters of other areas show a greater variation in water types. Despite these variations, chloride is by far the major anion in all water analyses from these areas .

Mink (1961) pointed out that cation exchange is a very dominant process in the ground waters of the southern Oahu basal-water area. As shown in figure 1, the cation-exchange process in the Pearl Harbor area exhibits an interesting trend with depth. This trend is toward increased per-

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centages of calcium and magnesium down to a I certain depth in the aquifer, where the trend re-verses and the percentage of sodium then increases with depth.

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Maximum and Minimum Chloride Levels in the Volcanic Bedrock Aquifer. Except where retarded by dikes, caprock, or artesian pressure, sea water freely intrudes Oahu's permeable lavas.
The rate of sea-water intrusion is generally greatest in areas of significant ground-water develop-The rate of sea-water intrusion is generally greatest in areas of significant ground-water development and heavy pumping.

Ground water is heavily developed from

basal-water sources in the Honolulu, Pearl

Harbor, and Waialua areas. Until the end of 1971 basal-water sources in the Honolulu, Pearl when sugarcane cultivation ceased, it was also heavily developed in the Kahuku area. Dikeimpounded water is developed in windward Oahu, in the Schofield area, and in the Waianae area. Draft is steady in the Honolulu area during the I

normally wet months of October to March each year, increasing by about a third during the rest of the normally dry months . Draft in the Pearl Harbor area is heavy from about April to October each year during the sugarcane-irrigation period. The draft decreases by about 75 percent during the wet months of October to March, when sugarcane is not irrigated. In the predominantly agricultural areas of Kahuku (before 1972) and of Waialua, draft is reduced by more than 95 percent during the nonirrigation period.

Plate 5 shows minimum chloride levels for both the high-level and basal-water aquifers; mostly for the 1968-70 period. In a few instances where information was not available for this period, earlier data are included.

Plate 6 was prepared to depict areas where a significant change in chloride content of the water occurs as a result of sea-water intrusion induced by pumpage. This map shows the maximum

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chloride-content levels for the same sources and same period as shown on the map for the minimum chloride-content levels (plate S) . Inland movement of intruding sea water is shown by the changing position of the isochlors on the two maps. However, even though these data show great increases in chlorinity as a result of pumpage in agricultural areas , during periods of nonirrigation the water of the aquifers will be somewhat freshened on an annual cycle by recharge from rainfall.

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Chloride Content by Depth Within the Volcanic Aquifer. The maximum and minimum chlorinity maps indicate that sea water intrudes laterally inland. However, water quality in a Ghyben-Herzberg lens is gradational from fresh to salt water in a vertical direction through a zone of transition. Thus, a knowledge of sampling depths is fundamental to understanding the mechanics and degree of sea-water intrusion in the artesian basal-water system. Although limited in precision and quantity, more depth information is available for Oahu than for the other islands. The data, where available, show where pumping has caused upconing and mixing of the higher-chloride waters in the areas of draft.

The map of plate 7 shows lines of equal depth to the 250-mg/1 chloride-content level, developed from depth-salinity relations of water from deep wells. Where sufficient data are available, depth-salinity curves are shown for selected one-minute rectangles . The lines of equal depth and the depth-salinity curves show generally the areas where the fresh-water lens is thickest and where it has been reduced by upward movement of the transition zone.

The contours and curves shown on plate 7 are preliminary and are, in some areas, prepared from very limited data because, heretofore, depth-oriented sampling has been little stressed. The preliminary data may, however, facilitate planning of a network of observation wells needed to better define movement of the salt-water body into the aquifers.

As shown, the basal water beneath Honolulu has less than 250-mg/1 chloride, and is more than 1,000 feet (305 meters) thick ln areas where the water level is 28 feet (8.5 meters) or more above mean sea level. The effects of mixing by pumping and tidal fluctuation, causing upward movement of the salt-water body, are quite apparent in the Pearl Harbor area. Even though the water-level elevation is only about 10 feet (3. 0 meters) less . than that in Honolulu, water fresher than 250 mg/lis generally less than 250 feet (76 meters) . thick. In the southern part of Pearl Harbor, and $\frac{1}{2}$ in areas of Kahuku, the entire ground-water body has a chlorinity greater than 250 mg/1.

The depth-salinity curve for the oneminute rectangle identified by 3406 in north-

central Oahu shows a significant decrease in chlorinity with depth in the' upper part of the basal lens. This phenomenon, although not as marked as in north-central Oahu, is also common in the Pearl Harbor and Kahuku areas. It results from irrigation with water higher in chlorinity than the underlying ground-water body. Infiltration of this water creates a reverse salinity gradient above the normal salinity gradient from underlying saline water.

Nitrate Levels. Nitrate concentrations above $1.1 \text{ mg}/1$ indicate other than natural conditions. Study of the levels of nitrate in ground waters is, therefore, useful to delineate areas where man has affected water quality. Although sewage effluent may have a localized effect, the major source of widespread and significant levels of nitrate on Oahu is fertilizers in irrigation return water.

Figure 2 shows areas of Oahu where information on nitrate levels is available. The map shows that the areas of lowest nitrate concentrations exist where man has had little effect upon the landscape, such as in the steep Koolau Range. The areas of greatest nitrate concentrations correlate well with known areas of irrigation.

Nitrate concentrations are greatest at the land surface and decrease with depth below irrigated areas. Thus, the values shown on the map, which were from greatly varied depths, show only areas of relative concentration and the exact concentrations within any area may have great variation.

Although the "greater than 8.0 mg/1" class . does not indicate an upper limit to the nitrate concentration, nearly all samples within this class range from 8 to 10 mg/1. However, two samples from shallow depths have high enough nitrate concentrations to warrant specific mention. One sample with a concentration of 19 mg/lwas from calcareous deposits in the Ewa plain of southwestern Oahu. The other sample, also from calcareous deposits, was from a well in Waimanalo. This sample had a nitrate concentration of 26 mg/1.

Calcareous Deposits in Kahuku Area. As mentioned previously, the fundamental factors controlling the chemical character of basal waters in the Kahuku area are high draft rates and cation exchange. However, extensive exposures of calcareous deposits in upland and recharge areas also directly affect the quality of fresh water in the basal lens. This is shown by the ratios of calcium to chloride for those waters with less than 250 -mg/l chloride, and by comparing the Kahuku values with other areas lacking these deposits in their recharge areas. Figure 3 shows that the ratios in the Kahuku area are significantly greater than in other areas.

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Figure 2. Maximum nitrate values for ground waters, 1962-72.

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Molokai, smallest of the five major islands; is composed of two volcanic mountains interconnected by a central plateau. The lavas are principally basaltic with areas of eastern and central Molokai overlain by andesite and even some trachyte (Stearns and Macdonald, 1947) . Annual rainfall has great areal variability, so that western and central Molokai receive from 15-30 inches (38-76 centimeters) annually, while eastern Molokai receives from 30-250 inches (76-635 centimeters) .each year, depending upon elevation. As determined by geologic features, ground water occurs as perched, dike-impounded, or basal water.

Ground-Water Zones. Perched water occurs primarily on the western slope of eastern Molokai. Here, ground water is perched by nearly impermeable ash flows within the permeable volcanics. Discharge from springs and tunnels tapping this zone fluctuates greatly. Unfortunately, only limited chemical data are available for these water sources .

Dike-impounded waters occur principally in a zone approximately 2 miles (3.2 kilometers) wide located in the higher·elevation areas of eastern Molokai, as shown on plate 8. In this area, the nearly impermeable, almost vertical dikes form a honeycomb structure of subsurface reservoirs. At the peak of the mountain formed by this resistant rock unit, rainfall is greatest of anywhere on the island.

The basal-water zone underlies almost the entire island, except for possibly beneath the rift zones. Because of greater rainfall, the freshwater lens of eastern Molokai has developed a far greater thickness than in the western or central parts of the island. Low-permeability sedimentary caprock retards flow of fresh ground water to the sea in the southern parts of eastern Molokai, causing further thickening of the lens.

Numerous shallow drilled wells and shafts acquire water from the volcanic basal aquifer in the southern coastal area of eastern Molokai, and many shallow dug wells tap the sedimentary caprock for a water supply.

Water Quality by Area. Only seven complete chemical analyses of ground water on Molokai were available for this study. The water sampled was from varied sources over a great period of time, which makes it difficult to show meaningful areal or change-with-time comparisons among the data. However, numerous and more current chloride data supply a large amount of information for areas deficient in more detailed analyses.

-In general, the ground water of the eastern

section of the island is of good quality and is only slightly affected by sea-water intrusion in the southern basal aquifer. In contrast, the water of the western and central parts of Molokai shows a large increase in chloride and dissolved-solids concentration over that of the south, with a predominant influence of intruded sea water (see plate 8).

The one complete analysis of dike zone water shown on plate 8 is a sodium calcium bicarbonate type water. This water has a similar chemical composition to the dike waters of the other islands. One distinct difference is the sulfate concentration. which is three to six times greater than in other dike waters. This may be the result of a reducing environment in the swamps upgradient from the well, but the exact origin is not known.

The basal water in the volcanic and sedimentary deposits in the southern coastal area of eastern Molokai has a relatively low chloride concentration. In addition, it does not exceed drinking water limitations for any chemical parameters measured. Waters in the volcanic aquifer from shafts 4 and 6 are sodium magnesium chloride bicarbonate and sodium chloride types, respectively. The variations in dissolved-solids concentrations and water-quality types for these sources may be a result of greater recharge to shaft 4 from streamflow infiltration from Kawela Stream.

The one chemical analysis from a shallow^{*}dug well in the sedimentary aquifer shows a dissolvedsolids concentration similar to the water from shaft 4. The water from the dug well, however, is a calcium magnesium bicarbonate type.

The great increase of chloride concentrations from less than 100 mg/1 in eastern Molokai to greater than 2,800 mg/l in western Molokai vividly shows the thinness of the fresh-water lens and the greater. saline-water intrusion in the western and central sections of the island.

The three chemical analyses for western and central Molokai show the water to be of two types. They are either sodium magnesium chloride or magnesium calcium chloride types. Of these sources, only the water from well 16 has chloride and dissolved-solids concentration within the limits for drinking water.

The analyses indicate that much of the intruded saline water beneath Molokai is chemically different from sea water. Comparisons of the analyses of samples from wells 1 and 15 with that of sea water diluted *to* the same respective concentrations appear to indicate this difference 追溯 due to cation exchange (see table 2). Worthy of \mathbb{R} notice is the fact that the temperature of the water. from well 1 was 93^{C} (34.0 C) when sampled. This suggests that some dissolved constituents may have been brought upward by volatiles or leached from rocks by rising gases (Stearns and Macdon 1947, p. 78).

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Table 2. Comparison of major cation and anion concentrations of water from well sources on Molokai with those of sea water diluted to the chloride content of the well water

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Maul, second largest of the islands, had its origin from two major volcanoes, Haleakala on the east and West Maui. The isthmus in the central portion of the island developed from the lava from · Haleakala banking against West Maul Mountain (Stearns and Macdonald, 1942) .

The principal aquifer is the volcanic rock, generally the highly permeable primitive basaltic lava flows , which make up the bulk of the island . In the isthmus and along the northern and southwestern shores of eastern Maui, the surface rocks are mainly calcareous dune sand. The sand is fairly permeable, and numerous wells have been drilled into it for home-lawn irrigation and for disposal of storm runoff. Consolidated sedimentary rocks are unimportant as aquifers because they are poorly permeable and generally lie above the water table.

Development and use of ground water and

the importation of surface waters, especially for sugarcane irrigation, have affected the quality of the basal water in many areas of Maul. In general, the quality of basal water will be degraded where natural recharge has been decreased by development, and the quality will be improved where natural recharge has been increased. Streamflow diversions at high altitudes on the northern flanks of Haleakala and in the northeastern and northwestern parts of western Maui have significantly reduced natural recharqe to the basal-water body in these areas. The recipient areas for the diverted water are those sections of Maui where sugarcane is irrigated. Here, recharge to the groundwater body is increased over natural conditions by the infiltrating irrigation water. Where streams are diverted at higher elevation and used for sugarcane irrigation at a lower elevation within the same drainage basin, net recharge to ground water may be reduced because of increased evapotranspiration.

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Figure 4. Maximum bicarbonate concentrations and generalized geology, 1962-72.

Fiqure 5. Maximum nitrate concentrations for qround-water sources, 1962-72.

Water-Quality Types. The generalized waterquality types for Maul's ground-water sources are depicted on plate 9. In cases where a significant variation in quality exists for a single source at different times, both water-quality types are shown.

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Only one analysis of dike-impounded water is available for Maui. The sodium bicarbonate type water from this source (5332-02) is of excellent quality. The analysis showed very low concentrations of all chemical parameters.

The remaining samples are believed to be from the basal-water system. The water is generally of a lower quality than the dike water. In many cases chloride and dissolved-solids concentration exceeds the limits for drinking water. The water is predominantly a sodium chloride type, which results from the high permeability of the volcanic rock aquifers throughout the island These highly permeable aquifers, in free communication with the sea, allow the development of only a shallow fresh-water lens and permit extensive sea-water

intrusion. Consequently, even thouqh they are diluted by fresh water in some areas, the qround waters throughout most of the island are quite similar to sea water in chemical percentages. The high-pumpage rates for sugarcane irrigation from March to late summer further accelerates seawater intrusion. This results in increased concentrations of all parameters associated with sea water, especially chloride. However, recharge from the imported surface water, and from rainfall during the nonpumping season freshens the water of the wells on a seasonal basis. Well 5321-0lis a good example of this freshening characteristic.

The few analyses of ground water along. the southern and southwestern margins of western Maui that are sodium chloride types, appear to have their own distinctive characteristics. These waters contain much higher percentages of calcium and magnesium than other ground-water samples from Maui. A further characteristic common to each of these analyses is that each of them is from

a source either beneath or inland of an extensive alluvial deposit. As on other islands, this chemical condition can be interpreted as the result of cation exchange. Because of the less extensive alluvial deposits on Maui, cation exchange is not as prominent as on Oahu.

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> One water-quality characteristic of particular significance on Maui is the high bicarbonate concentration shown in many ground-water analyses. The distribution of these high values is shown in figure 4. The concentrations greatly exceed the 140 mg/1 bicarbonate concentration of pure sea water and are significantly higher than anywhere else in the State. These very high concentrations may be the result of continuous recycling of irrigation water through the vegetation root zone. However, the extensive calcareous deposits of the isthmus may be a major contributing factor.

> Chloride Content of Ground Waters. The . magnitude of sea-water intrusion induced by pumping for irrigation is reflected in the chloride concentrations shown on plate 10. Because high rates of pumping are maintained during the spring and

summer months, these chloride concentrations may increase more than tenfold on an annual basis. The lines of equal chloride shown on the map represent low-chloride conditions during the winter months when little or no pumping occurs. The annual range in chlorinity from the winter to summer months is shown for the heavily pumped wells. The lines of equal chloride in the Kihei-Makena area outside the irrigated area of eastern Maui represents average conditions. Water-level contours in eastern Maui were drawn from measurements taken after 10 days of nearly complete pump shutdown. The contours, as drawn in the area of the isthmus, infer by their shape that flow of ground water is from the east and the west, and reflects, in part, the effect of recharge to the basal-water body by irrigation water return.

Nitrate Levels. The concentration of nitrate in the ground water of Maui is shown in figure 5, and is useful in delineating the extent of irrigation return water to the basal-water system. The map shows that no source exceeds the recommended domestic limit for nitrate of 45 mg/1,.

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Hawaii, the largest and southernmost of the islands, has an area of about 4, 030 square miles (10, 438 square kilometers). It was formed by the building and coalescence of five volcanoes: Kohala Mountains have long been extinct; Mauna Kea has had Holocene but no historic eruptions; Hualalai last erupted in 1801; Mauna Loa last erupted in 1950; and Kilauea is active today. The peaks of these volcanoes range in altitude from a few thousand to almost 13,800 feet (4,206 meters), resulting in about 80 percent of the island being above 1, 000 feet (305 meters) in altitude.

The bulk of the island is composed of moderately to highly permeable basaltic lava flows, which comprise the principal aquifer, These basaltic flows are capped by andesite on much of Mauna Kea and by andesite and trachyte flows on part of Kohala Mountains (Stearns and Macdonald, 1946) , Sedimentary deposits are generally thin and sparse.

Rainfall ranges from 10 to 300 inches (25-760 centimeters) annually, depending upon location and altitude. Island-wide average annual rainfall is 50 inches $(130$ centimeters), and because of the high permeability of the lavas a large part of the rainfall infiltrates to the water table. Some ground water is perched in the lava flows above sea level and some is impounded by dikes in rift zones . Most of the ground water, however, exists in the basalwater system. Unless evaporated, transpired, or intercepted for man's needs, it escapes to the sea as diffuse flow at or near the shore (Davis and Yamanaga, 1973).

Water-Quality Types. Of the water-quality analyses for the island of Hawaii depicted on plate 11, two are from samples of perched water. The water from tunnels 6 and 7 in Kohala Mountains is believed to be perched upon soil or weathered zones within the volcanics. They are sodium calcium bicarbonate and sodium maqnesium bicarbonate waters of excellent quality, with low concentrations of all chemical parameters.

The numerous other analyses on plate 11 are of basal-water samples from sources within. the permeable volcanics of the island. A great majority of these waters are sodium chloride types. In many of these analyses, the chloride and dissolvedsolids concentration exceeds the limitations for domestic drinking water. These conditions are the result of widespread sea-water intrusion into the highly permeable volcanic rock aquifer. Water levels are frequently less than 5 feet (1. 5 meters) above msl in the basal-water system, and there is no thick fresh-water lens. The areal extent of the sea-water effects in any specific area can be seen on plate 11 by comparing the circle sizes (total concentration of major ions) with their distance from shore; the lesser sea-water effect

having the smaller circle.

Because the effect of sea-water intrusion is minimal in the basal-water sources farthest from the shore, water samples from these areas have greatly varied chemical compositions. The inland wells appear to tap only the uppermost layer of... the fresh-water lens as recharged by infiltration. The waters from these wells have less than 150 · mg/1 dissolved-solids concentration, and the dominant anion is bicarbonate. The major cations may vary between different areas and within the same locale. In the Keauhou and Pahoa areas. these low-concentration waters are of sodium bicarbonate type but near Pepeekeo, the waters are of magnesium calcium bicarbonate type. However, in other areas, such as near H1lo and Honuapo, the cations vary greatly in percentages and no one cation is dominant for each place at all times .

Effects of Recent Volcanic Activity. Recent volcanic activity on the island of Hawaii appears to affect some ground water in and near the activity sites. One water analysis which exhibits these \mathbb{Z} effects is from a well within Kapoho cone $(3080-01)$ where the water is a calcium magnesium sodium bicarbonate type. The dissolved-solids concentration range is from 498 to 895 mg/1. The high bicar- $\sqrt{2}$ bonate concentration, which is believed to be indicative of volcanic gases at this location, ranges from 273 to 976 mg/l for three samples taken in 1968, 1970, and 1972.

Another well which appears to exhibit effects of volcanism is well 4858-01 in the Hualalai Volcanics. The sodium chloride type water from this well also has a relatively high bicarbonate concentration. The range of bicarbonate for three samples taken in 1960, 1965, and 1972 is from 260 to 439 mg/1. The possibility of volcanic gases entering this water source is further supported by fluoride concentrations of 2.0 and 2.2 mg/1 from two analyses made in 1965 and 1972 by two different agencies. These fluoride concentrations are much greater than those of other water in Hawaii, which rarely exceed 0.4 mg/1 of fluoride.

Chloride Levels. Because of the variations in recharge, pumping, and tidal effects within the permeable aquifers, the fresh water is readily mixed with the sea water. This results in brackish water at the coastline, even in. some areas where ground-water flow is large. In dry areas where ground-water flow is small, the basal water may be brackish for more than a mile (1.6 kilometers) inland. Because of this, sodium chloride waters dominate the coastal areas, where chloride concentrations of the basal waters clearly show the inland extent of sea-water intrusion. The chloride contents of the basal-water system on Hawaii are shown on plate 12. Levels of equal chloride content are drawn wherever adequate data are available.

e from U.S. Geological Survey 1+125,000 topographia map, 1944

Compiled by L.A. Swein, 1973

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Conclusions

From collection and evaluation of all available data, the first synthesis of qround-water quality for the five major islands of Hawaii is now available.

In qeneral, the water quality of each island is a function of the environment and of man's activities . The environmental factors are rainfall, geology, and sea-water intrusion. Man's activities that affect quality are aqricultural pumpinq and return irriqation water, industrial wastes, and domestic wastes.

Heavy pumpinq siqnificantly increases chloride concentrations by inducing sea-water intrusion. Fertilizer application adds soluble products to the soil, which are washed down with infiltratinq irriqation and rainwater. Nitrate is the most definitive chemical identifiable from irrigation return water. No current detrimental effects of industrial wastes upon the qroundwater bodies have been measured. Only limited information reqardinq the effects of solid and liquid domestic wastes is available. The principal or dominant factors affectinq each island may vary qreatly.

Kauai's ground-water quality is mainly a result of geoloqic effects . The few analyses available indicate calcium bicarbonate and magnesium sodium bicarbonate types of perched water. The dike-impounded water is also a magnesium sodium bicarbonate type water. Both of these types of water have low-chloride and dissolved-solids concentrations. The basal waters have higher concentrations of all chemical parameters. Bicarbonate is the principal major anion in the basal-water resource. The cations in these sources vary considerably. In a few analyses , the major anion is chloride. In these waters there appears to be significant cation exchange.

Water from Oahu's calcareous sedimentary aquifers is either a sodium chloride or sodium magnesium chloride type. All samples from calcareous materials have concentrations of chloride exceeding drinking water limits. In addition, all samples have bicarbonate concentrations greatly exceeding that of fresh or sea water, indicating solution of the calcareous rocks .

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High-level water on Oahu is primarily a sodium magnesium bicarbonate type with lowchloride and dissolved-solids concentrations. Exceptions exist in coastal areas of both mountain ranges and a qeothermal area in the Waianae Range.

Basal water in the three subzones of northcentral Oahu is distinctive for each subzone. De-

pending upon its location, the water is either a sodium magnesium chloride type or a sodium chloride type with varying percentaqes of ions. Basal water in the Kahuku area is a sodium maqnesium chloride type in the southern half of the area. In the northern half, where calcareous deposits and large draft are present, the. water is a magnesium calcium chloride type. Basal-water quality of southern Oahu varies greatly but is primarily a sodium chloride type. There is significant cation exchange in the deeper, more concentrated water sampled.

Areal rainfall variation is a definite factor determining Molokai's qround-water quality. Water in eastern Molokai's dike zone is a sodium calcium bicarbonate type with a higher sulfate concentration than that of other island dike waters . The two basal-water samples in the volcanic aquifer of eastern Molokai are sodium magnesium chloride bicarbonate and sodium chloride types. One shallow dug-well sample in the noncalcareous sedimentary aquifer is a calcium magnesium bicarbonate type. There is a great increase in salinity of ground waters from eastern Molokai to western Molokai. Samples from western and central Molokai are sodium magnesium chloride or magnesium calcium chloride types. The chloride concentration in one sample from the westernmost source was 2, 800 mg /1.

Dike water in Maui is a sodium bicarbonate type of low concentration for all parameters. Most basal-water samples on Maul are sodium chloride types as a result of sea-water intrusion. Where alluvial cap rock exists, there is a cation exchange effect upon the intruded waters. In much of Maul's ground water, there are very high concentrations of bicarbonate compared to natural water. This results from continuous recyclinq of irrigation water through the vegetation root zone, and also from dissolution of the calcium carbonate rock through which it infiltrates. Nitrate concentrations show the areas of greatest amount of return irrigation water.

Perched water sources on the island of Hawaii are sodium calcium bicarbonate or sodium magnesium bicarbonate types. The basal-water samples are mostly sodium chloride types The dissolvedsolids and chloride concentration often exceeds domestic usaqe limitations as a result of widespread sea-water intrusion. However, some areas inland and in the upper portion of the lens have good quality water with low-dissolved solids. The effects of recent volcanic activity are evident in two ground-water samples.

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Need for Further Investigations

Although this study has established general . baseline data for Hawaii's ground-water quality, it has also brought to light the many deficiencies in the data base that must be resolved in order to effectively understand and monitor changes in water quality.

Intensive study of the uncontaminated waters of high-level areas is severely limited by the scarcity of rainfall chemistry data. There is little information for the islands, which shows the effects of man's activities upon this rainfall chemistry. In addition, because of limited data, it cannot be determined what part of the chloride concentration in ground water is from rainfall and what part is from accumulated "dry rain" being washed from the surface and carried into the ground water.

To understand the variations in basal-water quality, the information must include depth from which the sample is taken. This depth information will help in the determination of the stratification of chloride concentration and its movement. In addition, depth information for complete chemical analyses will make it possible to determine where cation exchanqe is the most intense, and how water quality varies within aquifers. It is no longer acceptable to blindly compare analyses areally, when one sample may be from 20 feet (6.1 meters) below msl and another may be from 1, 000 feet (305 meters) below msl.

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If we are to adequately understand the ground-water systems and prevent contamination of our ground waters, measurements must be made for some parameters about which we presently have almost no information for the islands. Such

parameters would include biochemical oxygen demand, hydrocarbons, minor metals, and organic compounds.

It would greatly enhance the value of all chemical analyses if pH and temperature were determined at the sample site .. The prevalent practice of only determininq these parameters in the laboratory severely limits usefulness of the data. The most reliable possible information is needed for the study and interpretations of chemical reaction which take place in the ground-water bodies.

This report only touches upon the numerous chemical characteristics and anomalies present in Hawaii's ground-water bodies. Far more intensive study of the water quality of each area, and the anomalies, will be necessary to determine the origin and causes of each specific water-quality condition.

The historical and current data obtained for this study will establish a broad water-quality data base for today's conditions. However, water quality is continuously undergoing change, and as waste disposal on land and by subsurface injection becomes more widespread, the effects upon future water quality may be significant. It is inevitable that Oahu's ground-water quality will undergo change as urban sprawl spreads from the "protective" caprock area of southern Oahu out over recharge areas for ground-water supplies. Thus, priorities and methods to meaningfully monitor change must be established if undesirable contamination of our present and future water supplies is to be prevented.

REFERENCES CITED

.,,

Bryson, L.T., 1951, Chemical characteristics of Honolulu artesian waters: Honolulu Board of Water Supply, unpublished report, 47 p.

Davis, D.A., and Yamanaga, George, 1973, Water resources summary, island of Hawaii: Hawaii Div. Water and Land Devel. Rept. R47, 42 p.

Davis, S.N., 1969, Silica in streams and ground water of Hawaii: Univ. of Hawaii, Water Resources Research Center Tech. Rept. 29.

Eriksson, Erik, 1957, Thechemical composition of Hawaiian rainfall: Tellus, v. 9, p. 509-520.

Hem, J.D., 1970, Study and interpretation of the chemical characteristics of natural water: U.S. Geol. Survey Water-Supply Paper 1473, Second ed., 363 p.

Honolulu Water Commission, 1917, Untitled final report of the Commission, unpublished: City and County of Honolulu, v. IV and V.

Hubbard, M.K., 1940, The theory of qround-water motion: Jour. Geol., v. 68, no. 8, p. 924-926.

Macdonald, G.A., and Abbott, A.T., 1970, Volcanoes in the sea, The geology of Hawaii: Honolulu, Hawaii, Univ. of Hawaii Press, 441 p.

Macdonald, G.A., Davis, D.A., and Cox, D.C., 1960, Geoloqy and qround water resources of the island of Kauai, Hawaii: Hawaii Div. Hydrography Bull. 13, 212p.

Mink, J .F., 1961, Some geochemical aspects of sea-water intrusion in an island aquifer: Internat. Assoc. Sci. Hydrology Comm. of Subterranean Waters Pub. 52, p. 424-439.

 $_$ 1962, Excessive irrigation and the soils and ground water of Oahu, Hawaii: Science, v. 135, no. 3504, p. 672-673.

Rainwater, F .N., and Thatcher, L .L., 1960, Methods for collection and analysis of water samples: U.S. Geol. Survey Water-Supply Paper 1454, 301 p.

Rosenau, J .C., Lubke, E.R., and Nakahara, R.H., 1971, Water resources of north-central Oahu, Hawaii: U.S. Gaol. Survey Water-Supply Paper 1899-D, 39 p.

Stearns, H.T., and Macdonald, G.A., 1942, Geology and ground-water resources of the island of Maul, Hawaii: Hawaii Div. Hydrography Bull. 7, 344p.

_____ 1946, Geoloqy and ground-water resources of the island of Hawaii: Hawaii Div, Hydrography Bull. 9, 344 p.

____ 1947, Geology and qround-water resources of the island of Moloka1, Hawaii: Hawaii Div. Hydrography Bull. 11, 113 p.

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Stearns, H.T., and Vaksvik, K.N., 1935, Geology and ground-water resources of the island of Oahu, Hawaii: Hawaii Div. Hydrography Bull. 1, 479 p.

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- Takasaki, K.J., Hydrologic conditions related to subsurface and surface disposal of wastes in Hawaii: U.S. Geol. Survey open-file report, Water Resources Inv. 1-74, 5 sheets.
- Takasaki, K.J., and Valenciano, Santos, 1969, Water in the Kahuku area, Oahu, Hawaii: U.S. Geol. Survey Water-Supply Paper 1874, 59 p.
- Tenorio, P.A., Young, R.H.F., and Whitehead, H.C., 1969, Identification of irrigation return water in the sub-surface, Phase II: Univ. of Hawaii, Water Resources Research Center Tech. Rept. 33, 90 p.
- Tenorio, P.A., Young, R.H.F., Burbank, N.C., Jr., and Lau, L.S., 1970, Identification of irrigation return water in the subsurface, Phase III, Kahuku, Oahu and Kahului and Lahaina, Maul: Univ. of Hawaii, Water Resources Research Center Tech. Rept. 44, 53 p.
- U.S. Public Health Service, 1962, Drinking water standards, 1962: U.S. Public Health Service Pub. 956, 61 p.
- Visher, F.N., and Mink, J.F., 1964, Ground-water resources in southern Oahu, Hawaii: U.S. Geol. Survey Water-Supply Paper 1778, 133 p.

Zaidi, S.I., 1973, Determination of organic contaminants in Oahu ground water through carbon chloroform extraction {personal communication) .

APPENDIX A

Selected Chemical Analyses of Ground Water

The headings and the data included in the table columns on the following pages are:

The analyses given in the following table are as reported by the analyzing laboratory, except that where the analyzing laboratory failed to report calculated solids, hardness (Ca, Mg), noncarbonate hardness as carbonate alkalinity, these values were calculated and added by the author. With this exception, the accuracy of the reported analysis is solely the responsibility of the reporting laboratory.

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CHEMICAL ANALYSES OF OAHU GROUND WATER

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CHEMICAL ANALYSES OF MOLOKAI GROUND WATER

CHEMICAL ANALYSES OF MAUI GROUND WATER

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