

1 **An improved algorithm for unmixing first-order reversal curve diagrams using**
2 **principal component analysis**

3
4 Richard J. Harrison¹, Joy Muraszko¹, David Heslop², Ioan Lascu³, Adrian R. Muxworthy⁴, Andrew
5 P. Roberts²

6
7 ¹Department of Earth Sciences, University of Cambridge, Downing Street, Cambridge, U.K.

8 ²Research School of Earth Sciences, Australian National University, Canberra, ACT, Australia

9 ³Department of Mineral Sciences, National Museum of Natural History, Smithsonian Institution,
10 Washington, U.S.A.

11 ⁴Department of Earth Science and Engineering, Imperial College, London, U.K.

12
13 Corresponding author: Richard Harrison (rjh40@cam.ac.uk)

14
15 **Key Points:**

- 16 • The FORC-PCA unmixing method is tested on synthetic mixtures of magnetic domain
17 states.
- 18 • A new FORC-PCA algorithm is developed that solves the linear mixing problem for end
19 members with different domain states (including superparamagnetism).
- 20 • Feasibility metrics are developed that help to guide identification of physically realistic end
21 members.
22

23

24 **Abstract**

25 First-order reversal curve (FORC) diagrams of synthetic binary mixtures with single-domain, vortex
26 state, and multi-domain end members (EMs) were analyzed using principal component analysis
27 (FORC-PCA). Mixing proportions derived from FORC-PCA are shown to deviate systematically
28 from the known weight percent of EMs, which is caused by the lack of reversible magnetization
29 contributions to the FORC distribution. The error in the mixing proportions can be corrected by
30 applying PCA to the raw FORCs, rather than to the processed FORC diagram, thereby capturing
31 both reversible and irreversible contributions to the signal. Here we develop a new practical
32 implementation of the FORC-PCA method that enables quantitative unmixing to be performed
33 routinely on suites of FORC diagrams with up to four distinct EMs. The method provides access not
34 only to the processed FORC diagram of each EM, but also to reconstructed FORCs, which enables
35 objective criteria to be defined that aid identification of physically realistic EMs. We illustrate
36 FORC-PCA with examples of quantitative unmixing of magnetic components that will have
37 widespread applicability in paleomagnetism and environmental magnetism.

38

39 **1. Introduction**

40 Natural samples contain magnetic minerals with a wide range of grain sizes, domain states,
41 coercivity distributions, anisotropies, and interaction fields. First-order reversal curve (FORC)
42 diagrams provide a powerful method to characterize all these aspects of the magnetic mineralogy
43 (Pike et al., 1999; Roberts et al., 2000; Roberts et al., 2014), although their interpretation in the
44 literature is often based on qualitative assessments and empirical ‘fingerprinting’. Developments in
45 theoretical modelling (Muxworthy et al., 2004; Newell, 2005; Egli, 2006; Harrison & Lascu, 2014;
46 Roberts et al., 2017), new measurement protocols (Zhao et al., 2015, 2017), and new analysis
47 methods (Egli, 2013; Egli & Winklhofer, 2014; Heslop et al., 2014) have placed the processing and
48 interpretation of FORC diagrams onto a firm physical footing, which provides the opportunity for a
49 more quantitative approach to rock magnetic characterization. Application of principal component

50 analysis (Jolliffe 2002) to analyze entire sets of FORC diagrams (FORC-PCA) was introduced by
51 Lascu et al. (2015) as a quantitative method to unmix a suite of related samples into a linear
52 combination of up to four end members (EMs). The method has been applied successfully to unmix
53 the biogenic and detrital magnetic components of a sediment core from the Rockall Trough
54 (Channell et al., 2016), to characterize glacial/interglacial sedimentation on the Northwest Iberian
55 Margin (Plaza-Morlote et al., 2017), and to unmix the pedogenic and detrital magnetic components
56 of Minnesotan soils (Maxbauer et al., 2017). The key advantage of FORC-PCA lies in the two-
57 dimensional nature of the FORC diagram. Unmixing one-dimensional coercivity distributions (e.g.,
58 by fitting to the sum of standard basis functions) can be ambiguous, especially when there is strong
59 overlap between the coercivity distributions of different components (Heslop, 2015). The
60 information provided by the vertical B_u axis of a FORC diagram, however, provides additional
61 sensitivity to the presence of superparamagnetic (SP), single-domain (SD), vortex (V), and multi-
62 domain (MD) states, the ability to detect the presence or absence of interactions in each EM, and a
63 way to discriminate between minerals with different types of magnetocrystalline anisotropy. Note
64 that throughout this paper we follow Roberts et al. (2017) in referring to ‘vortex’ states rather than
65 to ‘pseudo-single domain’ or ‘PSD’ states. The term ‘vortex’ is broadly defined by Roberts et al.
66 (2017) to include both single-vortex (SV) and multi-vortex (MV) states, which more accurately
67 describe the physics of magnetic particles (Donnelly et al., 2017) in the intermediate size range
68 between the SD and MD states.

69
70 Despite the numerous advantages of FORC-PCA, the method has some shortcomings that currently
71 limit its usefulness for rock magnetic characterization. First, the FORC distribution is sensitive
72 primarily to the irreversible component of the magnetic response of a sample to a changing
73 magnetic field. Purely reversible responses to the changing field (e.g., from superparamagnetic,
74 paramagnetic, antiferromagnetic or diamagnetic sources), are strictly absent from the FORC
75 diagram, and cannot be identified as EM components in a FORC-PCA analysis. Even for

76 ferrimagnetic sources, the magnetization response of any sample can be split into the sum of
77 reversible and irreversible components, with the ratio of the two depending largely on the domain
78 state: SD states are dominated by irreversible magnetization, whereas MD states are dominated by
79 reversible magnetization. In its current form, therefore, the unmixing proportions reported by
80 FORC-PCA may deviate significantly from the actual proportions (by mass or volume) of the EMs
81 present in the sample, especially if the EMs represent populations of grains with different domain
82 states. Second, the process of choosing appropriate EMs (based often on a limited sampling of the
83 unmixing space by a dataset) can be subject to non-uniqueness, user subjectivity, and in the most
84 serious cases, to selection of physically unrealistic EMs.

85

86 In this paper, we develop an improved algorithm for FORC-PCA that addresses these issues. We
87 present a practical implementation that allows the FORC distribution and the FORCs themselves to
88 be reconstructed simultaneously, and describe objective criteria that can be used to guide the most
89 appropriate EM choice to enable quantitative unmixing of FORC diagrams. These improvements to
90 the FORC-PCA method are implemented and integrated into a new version of the FORCem
91 package within FORCinel (Harrison & Feinberg, 2008; Lascau et al., 2015), as described in the
92 supplemental material.

93

94 **2. Materials and Methods**

95 Three synthetic binary mixtures that contain known proportions of SD, V, and MD
96 magnetite were studied using FORC-PCA. Binary mixtures of SD-V and SD-MD particles were
97 kindly provided by Bruce Moskowitz of the Institute for Rock Magnetism. These samples have
98 been used in several previous studies of magnetic unmixing (Carter-Stiglitz et al., 2001; Dunlop &
99 Carter-Stiglitz, 2006; Lascau et al., 2010). The SD EM is a freeze-dried sample of a cultivated strain
100 of the MV1 magnetotactic bacterium. The magnetosomes have a well-constrained grain-size
101 distribution, with particle sizes of 35x35x53 nm aligned in chains of 10-20 crystals (Moskowitz et

102 al., 1993). Both the V and MD EMs are synthetic magnetites produced by Wright Industries, with
103 typical grain-size distributions of 1-3 μm and 8-40 μm , respectively (Carvallo & Muxworthy,
104 2006). The SD-V and SD-MD mixtures were produced by first dispersing the coarser EM in CaF_2 to
105 0.1% by weight and then adding MV1 to obtain the desired mass proportions. The V-MD mixture
106 was created by weighing the mass of EMs, and dispersing them to a 1% concentration in a fine-
107 grained sucrose matrix. Samples were mixed gently, placed into gelatin capsules and packed with
108 quartz wool or Kimwipe tissues to prevent vibration of the sample during measurement. The mixing
109 proportions of all samples are listed in Table 1. Bulk hysteresis parameters (newly measured for this
110 study) for all samples are listed in Table 2.

111

112 Measurements were made at the University of Cambridge on a Lakeshore PMC MicroMag
113 vibrating sample magnetometer. For each sample 174 FORCs were acquired in 1.5 mT field
114 increments with 200 ms averaging time. The FORC data were imported into FORCinel (Harrison &
115 Feinberg, 2008) and processed using the VARIFORC smoothing algorithm (Egli, 2013), resampled
116 on a 2 mT grid, and subjected to PCA analysis following the protocol described by Lascu et al.
117 (2015). EMs were selected from known pure samples to constrain the mixing space. In each case
118 the finest magnetic component was chosen as EM1. The synthetic binary mixtures were then
119 unmixed using PCA and the mixing proportions were calculated.

120

121 **3. Results**

122 Representative FORC diagrams for each set of studied binary mixtures are shown in Figs. 1-
123 3. The MD EM (Figs. 1a, 3c) has a typical MD FORC diagram dominated by a low-coercivity,
124 vertically spread signal and a weak, high-coercivity, horizontally spread tail, likely related to strong
125 pinning of domain walls by stress fields on surfaces and at internal defects/dislocations (Pike et al.,
126 2001a; Lindquist et al., 2015). The SD EM (Figs. 1c, 2c) has a typical SD FORC diagram for non-
127 interacting uniaxial SD particles, comprising an intense horizontal ridge and a corresponding weak

128 negative signal located close to the negative B_u axis (Muxworthy et al., 2004; Newell, 2005; Egli et
129 al., 2010). The coercivity distribution for the cultured MV1 bacteria is particularly narrow. Some
130 vertical spreading of the horizontal ridge is evident, which indicates either the presence of magnetic
131 interactions among the magnetosome chains, or a small degree of chain collapse (Li et al., 2012;
132 Egli & Winklhofer, 2014; Harrison & Lascu, 2014). A slight vertical offset of the horizontal ridge is
133 likely a viscous magnetization effect caused by the time asymmetry of the FORC measurement
134 protocol (Egli, 2013). The V EM (Figs. 2a, 3a) consists of an intense, closed-contour peak with
135 broad vertical and horizontal spreading, and three weaker, less prominent lobes. Although such
136 FORC signatures can be created by strongly interacting SD clusters (e.g., Carvallo et al., 2005), this
137 explanation can be ruled out here because the known grain size of the sample (1-3 μm) far exceeds
138 the upper SD threshold size. Instead, the broad central peak and three lobes are interpreted as MV
139 and SV processes, respectively, dominated by intra-particle, rather than inter-particle, interactions.
140 A weak negative signal close to the negative B_u axis is visible in Fig. 3a. For V-MD mixtures, the
141 maximum intensity of the FORC signal in each EM is comparable, so that both signals are clearly
142 evident in a ~50:50 mixture (Fig. 3b). The intense positive signal associated with the SD EM,
143 however, dominates the FORC signal of the SD-MD and SD-V mixtures (Fig. 1b, 2b), so that only
144 when >80% of the mixture is constituted by the MD or V EMs does their presence become obvious
145 in the FORC diagram.

146
147 Results of FORC-PCA analysis using the method of Lascu et al. (2015) are shown in Fig. 4. There
148 is a systematic non-linear deviation in all three mixtures between the FORC-PCA calculated (EM1)
149 and actual weight fractions (EM1*) of EMs used to prepare the samples. With EM1 defined to be
150 the finer-grained EM, all three binary mixtures have a concave down relationship between the
151 calculated versus actual mixing proportions. The non-linearity is most pronounced for the SD-MD
152 binary mixture, and is least pronounced for the SD-V binary mixture.

153

154 **4. Origin of the unmixing discrepancy**

155 The difference between mixing proportions derived by FORC-PCA and the known mass
 156 proportions of EMs in the synthetic mixtures (Fig. 4) can be explained by the fact that the SD, V,
 157 and MD states have different ratios of irreversible to reversible magnetization. The FORC
 158 distribution, ρ , is defined as:

$$160 \quad \rho = -\frac{1}{2} \frac{\partial^2 M}{\partial B_a \partial B_b}, \quad (1)$$

161
 162 where M is the magnetization, B_a is the reversal field and B_b is the measurement field. Prior to
 163 FORC-PCA analysis, each FORC diagram is normalized by its integral:

$$164 \quad \int \int \rho dB_a dB_b = M_s - M_{rev} = M_{irr}, \quad (2)$$

165
 166 where M_s is the saturation magnetization, M_{rev} is the reversible component, and M_{irr} is the
 167 irreversible component (Pike, 2003). FORC-PCA describes each normalized FORC diagram as the
 168 linear sum of normalized EMs, so that the mixing proportions are defined as (e.g., for a binary
 169 mixture):

$$170 \quad \text{EM1} = \frac{m_1 M_{s1} f_1}{M_{irr}}, \text{ and} \quad (3)$$

$$171 \quad \text{EM2} = 1 - \text{EM1} = \frac{m_2 M_{s2} f_2}{M_{irr}}, \quad (4)$$

172
 173 where m_1 and m_2 are the mass of each EM, M_{s1} and M_{s2} are the mass-normalized saturation
 174 magnetization of each EM, and f_1 and f_2 are the ratio of the irreversible magnetization to saturation
 175 magnetization for each EM. The mass proportions of the EMs in the mixture are given by:

179

180
$$EM1^* = \frac{m_1}{m_1+m_2}. \quad (5)$$

181

182 Rearranging equations 3 and 4 and substituting in 5, we obtain:

183

184
$$EM1^* = \frac{EM1}{EM1+f(1-EM1)}, \quad (6)$$

185

186 where

187

188
$$f = \frac{M_{s1}f_1}{M_{s2}f_2}. \quad (7)$$

189

190 The f factor expresses how different the EMs are in terms of their irreversible/reversible
 191 magnetization contributions. A value of $f = 1$ corresponds to ideal behavior ($EM^* = EM1$), and is
 192 obtained only when the EMs contain identical minerals ($M_{s1} = M_{s2}$) with equal ratios of irreversible
 193 to saturation magnetization ($f_1 = f_2$). Least-squares fits to plots of EM1 vs EM1* are shown as solid
 194 lines in Fig. 4, yielding f values of 2.72, 1.52, and 2.69 for the SD-MD, SD-V, and V-MD mixtures,
 195 respectively. All EMs contain magnetite; therefore, the different f factors indicate that the
 196 irreversible contribution to the magnetization of each domain state is different, with $SD > V > MD$.
 197 Although, in principle, equation 6 allows the FORC-PCA proportions to be corrected, prior
 198 knowledge of the f factor is required. This is not a practical solution when the properties of the EMs
 199 are unknown.

200

201 **5. An improved FORC-PCA algorithm**

202 The non-linear unmixing discrepancy documented above can be corrected by applying PCA
 203 to the FORC magnetization surface (which contains both reversible and irreversible contributions),
 204 rather than to the FORC distribution (which contains only irreversible contributions); the FORC
 205 magnetization surface has been shown previously to mix linearly (Muxworthy et al., 2005). This

206 approach, however, poses a challenge to interactive exploration of the unmixing space that is
 207 necessary to identify suitable EMs: as each point in the mixing space is explored, it becomes
 208 necessary to estimate ρ over the reconstructed magnetization surface to obtain the corresponding
 209 FORC diagram. Here we overcome this problem by applying PCA to the set of six polynomial
 210 coefficients that are used to fit the magnetization surface during smoothing of the input FORC
 211 diagrams (Pike et al., 1999). In this way, the reconstructed set of coefficients at any given point in
 212 the unmixing space can be used to calculate both the magnetization surface and its derivatives
 213 simultaneously.

214

215 Our procedure is described as follows. Raw FORC data for a set of samples to be analyzed are
 216 imported into FORCinel. A linear high-field slope correction is applied, and a record is kept of the
 217 mass normalized M_s value for each sample, for future reference. Here, the slope correction was
 218 performed by fitting a straight line to the high-field portion of the FORCs. In cases where the
 219 FORCs have not been measured to sufficiently high fields to fully saturate the ferrimagnetic
 220 component, it may be desirable to perform the correction using a separately determined value of the
 221 high-field susceptibility. The FORCs are normalized to $M_s = 1$, the lower branch subtracted
 222 (optionally), and processed using the VARIFORC variable smoothing algorithm (Egli, 2013). For
 223 consistency with the published code (see supplementary materials) we use the VARIFORC
 224 coordinate scheme (B_c, B_u) rather than the measurement coordinate scheme (B_a, B_b) in the
 225 following. For each output point in a processed FORC diagram, a weighted second-order
 226 polynomial fit is performed to the local magnetization surface over a rectangular area defined by the
 227 horizontal and vertical smoothing factors (Egli, 2013):

228

$$229 \quad M(B_c, B_u) = a_0 + a_1 B_c + a_2 B_u + a_3 B_c^2 + a_4 B_c B_u + a_5 B_u^2 \quad . \quad (8)$$

230

231 The FORC distribution (eqn. 1) is then given by:

232

$$233 \quad \rho = \frac{a_3 - a_5}{4}. \quad (9)$$

234

235 In order to analyze sets of FORC diagrams that may have been acquired using different
 236 measurement protocols, each polynomial coefficient in eqn. 8 is interpolated bi-linearly onto a
 237 rectangular grid, capturing a specified region of interest. For a rectangular grid containing N points,
 238 there will be $6N$ observations for each FORC diagram, corresponding to the six bi-linearly
 239 interpolated polynomial coefficients for each point. The FORC-PCA method of Lascu et al. (2015)
 240 is then applied, simply replacing the N values of the FORC distribution with the $6N$ polynomial
 241 coefficients for each sample. Once the number of significant PCs has been chosen ($n \leq 3$,
 242 corresponding to a maximum of 4 EMs), low-rank approximations of both the magnetization
 243 surface and the FORC distribution can be reconstructed for any chosen location within the resulting
 244 unmixing space (score plot). Exploring the unmixing space to identify potential EMs can now be
 245 performed interactively, guided by both the reconstructed magnetization and corresponding FORC
 246 diagram.

247

248 A complication occurs when the option to subtract the lower branch from the normalized FORCs
 249 prior to smoothing is chosen. Lower-branch subtraction was introduced by Egli (2013) to improve
 250 smoothing performance in the vicinity of the $B_b = 0$ axis (an axis extending from the origin at a -45°
 251 angle in VARIFORC coordinate space). Lower-branch subtraction reduces significantly the
 252 appearance of smoothing artefacts along this axis when using variable smoothing protocols because
 253 it removes sigmoidal magnetization contributions that are poorly described by a second-order
 254 polynomial. When smoothing is performed after lower-branch subtraction, the set of polynomial
 255 coefficients in eqn. 8 describes the subtracted magnetization surface rather than the desired full
 256 magnetization surface. In order to reconstruct the full magnetization surface, a second smoothing
 257 step is performed on a synthetic 2D magnetization surface created using just the lower-branch

258 signal. This lower-branch surface is fitted using eqn. 8, but with the strict constraint that $a_3 = a_5$,
259 thereby ensuring that the FORC distribution (and its associated artefacts) associated with the lower-
260 branch surface is zero (eqn. 9). Polynomial coefficients resulting from the fit to the lower-branch
261 surface are then added to those resulting from the fit to the lower-branch subtracted magnetization
262 surface, which are then used as input to the FORC-PCA. This double-smoothing procedure allows
263 the full magnetization surface to be reconstructed from the chosen PC combination, while retaining
264 an artefact-free representation of the reconstructed FORC diagram.

265

266 Heslop & Roberts (2012a) demonstrated that, because of the corrupting effects of measurement
267 noise, it is necessary to calculate statistical significance levels to identify the parts of a FORC
268 distribution where ρ is significantly above the signal-to-noise ratio. Use of PCA to provide a low-
269 rank approximation of a collection of measured FORC diagrams is also an effective approach to
270 reduce the influence of noise in representing a mixing system (Heslop, 2015). Therefore, while
271 PCA will not eliminate noise completely, its effect on the representation of the mixing system and
272 on the identified EMs is reduced substantially compared to individual FORC diagrams.

273

274 Results of the new algorithm applied to the synthetic binary mixtures are shown in Figs. 5-7. The
275 V-MD mixture (Fig. 5) is well described as a binary mixture, with 99% of the variance in the
276 dataset explained by PC1. Pure EMs are included within the dataset, which leads to no ambiguity in
277 the choice of EM1 (V) and EM2 (MD) (Fig. 5a, b). The SD-MD mixture (Fig. 6) can be
278 approximated as a binary mixture, with 95% of the variance being explained by the first principal
279 component (PC1). However, a small but significant second principal component (PC2) is needed to
280 bring the variance explained to >99% (Fig. 6g). Without including PC2, it is not possible to isolate
281 completely a pure MD EM. This effect is caused by subtle coercivity differences of the MV1
282 bacteria from sample to sample, which only become apparent because of the intense and narrow
283 nature of their FORC distribution. Possible explanations for the coercivity difference between

284 samples include differences in oxidation state that resulted from sample storage in air for over 10
285 years, or different degrees of bacterial chain collapse. By including PC2, small coercivity
286 differences can be taken into account, enabling a pure MD EM to be identified (EM1, Fig. 6a),
287 along with two SD EMs (EM2 and EM3) that differ only in their average coercivity (Figs. 6b and c;
288 Table 3). Hence, PC1 describes the binary mixing between SD and MD EMs, and PC2 accounts for
289 the varying coercivity of the SD MV1 component. A similar approach was taken to describe the
290 SD-V mixture (Fig. 7), although the coercivity variation of the MV1 samples is less pronounced
291 (99% of the variance is explained by PC1 alone). In all three cases, the mixing proportions derived
292 from FORC-PCA agree well with the known mass fractions. The 2σ differences between calculated
293 and observed proportions are 2%, 5%, and 6% for the SD-MV, SD-V, and MD-V binary mixtures,
294 respectively. These observations provide an empirical estimate of the error in the unmixing
295 proportions that is likely to be achieved using FORC-PCA in optimal cases (i.e., where the mixing
296 space is well sampled by the dataset).

297

298 **6. Feasibility metrics**

299 An inherent part of the FORC-PCA method is the supervised exploration of the unmixing
300 space in order to identify appropriate EMs (Lascu et al., 2015). This process is only unambiguous
301 when the sample set includes examples of each EM that is being solved for (as is approximately the
302 case for the binary mixtures studied here). When the sampling of the unmixing space is incomplete,
303 however, the method relies heavily on the expertise of the user to identify (a) EMs that enclose the
304 entire set of sample scores (with the exception of outliers identified by residual analysis), (b) pure
305 EMs (i.e., that do not contain any residual contributions from the other EMs), and (c) EMs that are
306 physically realistic (i.e., the reconstructed FORC diagram for each EM corresponds to an achievable
307 FORC geometry based on knowledge of the magnetic mineralogy and the responses that can be
308 modelled physically) (Harrison & Lascu, 2014). With access to only the reconstructed FORC
309 diagram, identification of physically unrealistic regions of the unmixing space relies on subjective

310 criteria. The availability of reconstructed FORCs, however, provides objective information from
 311 which criteria can be defined to assess the physical feasibility of the corresponding FORC diagram.
 312 Following the approach of Heslop & Roberts (2012b), three criteria that can be applied to assess the
 313 feasibility of reconstructed FORCs are: (a) saturation (i.e., no FORC should exceed the normalized
 314 value of $M_s = 1$), (b) monotonicity (i.e., the first derivative of a FORC with respect to the
 315 measurement field should remain nonnegative), and (c) crossing (i.e., the first derivative of the
 316 magnetization surface with respect to the reversal field should remain positive, meaning that
 317 FORCs do not intersect each other). Each of these metrics can be used on their own, or in
 318 combination, to define the region of unmixing space that is physically realistic. The EMs should be
 319 contained entirely within that region.

320

321 We define three metrics for each of the feasibility criteria, which vary from 0 (completely
 322 unsatisfied) to 1 (completely satisfied):

323

$$324 \quad m_{saturation} = \frac{\sum |M_A|}{\sum |M|}, \quad (10)$$

325

326 where M_A is the subset of the magnetization, M , that satisfies the condition $|M| \leq 1$;

327

$$328 \quad m_{monotonicity} = \frac{\sum \left(\frac{dM}{dB_a} \right)_A}{\sum \left| \frac{dM}{dB_a} \right|}, \quad (11)$$

329

330 where $\left(\frac{dM}{dB_a} \right)_A$ is the subset of $\frac{dM}{dB_a}$ that satisfies the condition $\frac{dM}{dB_a} \geq 0$; and

331

$$332 \quad m_{crossing} = \frac{\sum \left(\frac{dM}{dB_b} \right)_A}{\sum \left| \frac{dM}{dB_b} \right|}, \quad (12)$$

333

334 where $\left(\frac{dM}{dB_b}\right)_A$ is the subset of $\frac{dM}{dB_b}$ that satisfies the condition $\frac{dM}{dB_b} \geq 0$.

335

336 The metrics can be combined into a single feasibility metric, m , by multiplying them together in any
 337 combination. By calculating m over a grid of points, contours of feasibility can be used to indicate
 338 the region of unmixing space where the criteria are satisfied fully ($m = 1$). In practice, some
 339 allowance is needed for the fact that we are dealing with a low-rank approximation to the data, that
 340 some non-monotonicity may be genuinely present (e.g., for SP grains), and that experimental noise
 341 can cause FORCs to cross as saturation is approached. This means that m values slightly less than 1
 342 should be allowable. Here we take $m > 0.99$ as a reasonable (although arbitrary) guideline of
 343 acceptability (Fig. 6g and 7g). Given the essentially binary nature of the mixtures, the placement of
 344 EM3 slightly outside the $m = 0.99$ contour in Figs. 6g and 7g has been done to maintain EM2 and
 345 EM3 at a constant value of PC1 away from EM1, whilst ensuring that all data points are contained
 346 within the mixing triangle.

347

348 A fourth metric, which should be used independently of the other three, describes the amount of
 349 negative signal in the processed FORC diagram. Given that negative regions are an intrinsic feature
 350 of many FORC diagrams, this metric is less stringent than the others (values significantly < 1 are
 351 acceptable). However, there are specific domain states that do not have intrinsically negative
 352 regions, or have only weakly negative regions, so evaluating this metric can be helpful to define the
 353 location of specific EMs. For example, inappropriate appearance of strong negative signals can be
 354 caused by over-subtraction of other EMs, which provides a good indication that EM selection has
 355 strayed too far from the data. The positivity metric is defined as:

356

$$357 \quad m_{positivity} = \frac{\sum |\rho_A|}{\sum |\rho|}, \quad (13)$$

358

359 where ρ_A is the subset of the FORC distribution, ρ , that satisfies the conditions $\rho \geq 0$. Steep drops
360 in $m_{\text{positivity}}$ may indicate that over-subtraction of other EMs is occurring.

361

362 **7. Example**

363 To illustrate the new FORC-PCA algorithm applied to natural mixtures of different domain
364 states, we analyzed a suite of greigite-bearing clays from Florindo et al. (2007), which were
365 deposited between 800 ka and 600 ka in the Tiber River coastal alluvial plain around Rome. A total
366 of 17 FORCs were measured, 14 of which contain magnetostatically interacting SD greigite mixed
367 with varying amounts of a SP/SD greigite. The other 3 samples contain the SP/SD signal only. The
368 latter samples were significantly less magnetic than the former, and have noisy processed FORC
369 diagrams. FORC data from these three samples were averaged to produce a single representative
370 example of the pure SP/SD component. This averaged FORC and the other 14 FORCs were then
371 analyzed using FORC-PCA (Fig. 8). Only two PCs are needed to explain over 90% of the variance
372 in the dataset, with a third PC bringing the variance explained to 98%. For illustrative purposes, we
373 use a two-PC model constructed from PC1 and PC3, which provides the most convenient projection
374 of the key mixing trends. Three EMs are identified. Key features of EM1 (Fig. 8a) are a negative
375 region close to the negative B_u axis (1), a second negative region that is elongated and steeply
376 angled down and to the right (2), and a kidney-shaped positive peak that is strongly offset in the
377 negative B_u direction and extends only slightly above the $B_u = 0$ axis (3). All three of these features
378 are diagnostic of relatively weakly interacting SD greigite grains (Roberts et al., 2011) with cubic
379 magnetocrystalline anisotropy (Harrison & Lasca, 2014). Key features of EM2 (Fig. 8b) are a
380 negative region close to the negative B_u axis (1) and a rounded positive peak that is offset in the
381 negative B_u direction and extends far above the $B_u = 0$ axis (2). Both features are diagnostic of
382 strongly interacting SD greigite (Roberts et al., 2011; Harrison & Lasca, 2014). Key features of
383 EM3 (Fig. 8c) are a low-coercivity ridge with maximum intensity at 0 mT (1) and an increasing
384 positive signal extending along the negative B_u axis (2). Feature 1 is characteristic of non-

385 interacting SD greigite particles with coercivities that have been reduced by thermal activation
386 (Pike et al., 2001b; Rowan & Roberts, 2006). Feature 2 is likely due to viscous SP behavior, which
387 leads to the negative initial slope of each FORC (arrow in Fig. 8f; Pike et al., 2001b). However,
388 given that the intensity of this feature continues to increase, even as reverse saturation is
389 approached, it is also likely to be partially an instrumental artefact.

390

391 The choice of EMs in this case has been guided by the following principles. First, EM3 is fixed by
392 inclusion of the pure SP/SD EM in the dataset. This sample plots to the far right of the unmixing
393 space (Fig. 8g), close to but within the guideline boundary of physical feasibility. Moving left, away
394 from EM3 in a direction parallel to the PC1 axis (which describes the largest mode of variability in
395 the dataset), yields a binary mixture of EM3 and a moderately-interacting SD greigite EM (EM1).
396 The most extreme left-hand data point lies close to the guideline boundary of physical feasibility.
397 However, the FORC diagram for this data point contains a trace residual of EM3. In order to obtain
398 a pure EM, one must move further to the left. The guideline boundary of physical feasibility places
399 a limit on how far to the left one can go before the reconstructed FORCs for EM1 become
400 physically unrealistic. We place EM1 at the $m = 0.99$ threshold, which yields a physically realistic
401 pure EM with no residual trace of EM3. The placement of EM2 is more difficult because it lies well
402 within the guideline region of physical feasibility. Here, the positivity index (eqn. 13) provides an
403 additional guideline (inset to Fig. 8g). A steep drop in $m_{\text{positivity}}$ is observed if EM2 is placed too far
404 along the positive PC2 axis, which is caused by over-subtraction of EM1 from the reconstructed
405 FORC diagram. If EM2 is placed too far along the positive PC1 axis then not all the data are
406 enclosed by the unmixing space. Combined, these two principles place important constraints on the
407 location of EM2, and produce a reconstructed FORC diagram with a recognizable geometry and
408 minimal residual traces of EM1 and EM3.

409

410 Having defined the unmixing space, the proportions of the three EMs can be determined (Fig. 8h).
411 Two distinct mixing trends can be identified in the data: a mixing between SP/SD and strongly
412 interacting SD greigite, and one between weakly and strongly interacting SD greigite. The first
413 mixing trend can be explained by grain growth of authigenic greigite from small, non-interacting
414 particles below the SP threshold size to larger, stable SD particles in closely packed clusters within
415 framboids with strong interactions (Rowan & Roberts, 2006). The second mixing trend can be
416 explained as a weakening of the interactions between SD greigite particles, driven by a lowering of
417 the packing fraction. A possible mechanism to explain this trend is the progressive replacement of
418 strongly interacting greigite framboids by thermodynamically stable, paramagnetic pyrite. This
419 process was recently identified by Ebert et al. (2018) using high-resolution magnetic force
420 microscopy imaging. This interpretation is consistent with the lack of a mixing trend between EM1
421 (SP/SD) and EM3 (weakly interacting SD), which cannot be achieved in this pyrite replacement
422 scenario without first going through the strongly interacting SD greigite EM.

423

424 **8. Discussion**

425 Unmixing the magnetic properties of rocks, sediments, and soils is a primary task in rock
426 magnetism. Numerous methods exist to tackle this problem (e.g., Robertson & France, 1994;
427 Kruiver et al., 2001; Dunlop, 2002a, 2002b; Egli, 2004a, 2004b, 2004c; Franke et al. 2007; Heslop
428 & Dillon, 2007; Lascu et al., 2010, 2015; Ludwig et al., 2013; Lagroix & Guyodo, 2017) as well as
429 an extensive toolbox of magnetic proxies that are designed to highlight specific magnetic
430 mineralogy variations in environmental contexts (Evans & Heller, 2003; Liu et al., 2012). No single
431 method is perfect for all cases, and usually a combination of methods is needed to unmix all
432 magnetic components contained within a material. In particular, preparatory studies performed at
433 high sampling resolution provide an efficient way to prescreen a dataset, and to identify samples
434 that are closest to potential EMs (e.g., EM3 in Fig. 8). The FORC-PCA method is ideally suited to
435 characterizing ferrimagnetic minerals, with an emphasis on discriminating populations of grains that

436 differ in domain state, coercivity distribution, anisotropy, and interaction field (i.e., aspects to which
437 FORC diagrams are particularly sensitive). Here we have resolved many of the outstanding issues
438 associated with the original FORC-PCA method of Lascu et al. (2015), including solution of the
439 linear mixing equation, the ability to identify SP EMs that are dominated by reversible
440 magnetizations, and reducing ambiguities in defining the unmixing space. Excellent agreement
441 between our calculated proportions for SD-MV and SD-V mixtures contrasts starkly with attempts
442 to unmix these samples using either linear or non-linear mixing in a Day plot (Day et al., 1977;
443 Dunlop & Carter-Stiglitz, 2006). Failure of the Day plot unmixing approach was explained by
444 Dunlop & Carter-Stiglitz (2006) as due to the squareness of hysteresis loops for MV1 bacteria,
445 which violates the linear assumption of the unmixing model (Dunlop, 2002a, 2002b). This
446 illustrates one of the key advantages of PCA, which makes no prior assumptions about the shape of
447 the EM signals (Heslop, 2015).

448
449 The need to use three EMs to describe binary SD-V and SD-MD mixtures highlights an important
450 underlying assumption of the FORC-PCA method, namely that the properties of each EM are
451 constant throughout a sample set, with only the mixing proportions varying from sample to sample.
452 Whenever this assumption is not met, additional ‘fictive’ EMs may be needed to define adequately
453 the total variability within a dataset. This is clearly the case for the MV1-bearing mixtures, where
454 significant coercivity variations of the bacterial component exist from sample to sample. Given the
455 narrow coercivity distribution of the MV1 bacteria, use of a third EM becomes necessary to isolate
456 a pure V or MD EM. Most natural samples have broader coercivity distributions, however, and as
457 long as intra-EM variability is low compared to inter-EM variability, ‘fictive’ EMs are not typically
458 necessary. The likelihood that ‘fictive’ EMs will be needed to account for intra-EM variability
459 increases as the size of the FORC dataset increases. For large datasets, it may be necessary to
460 perform a series of FORC-PCA analyses on subsets of the data. This approach allows
461 commonalities between EMs extracted from different subsets to be identified, and the nature of

462 intra-EM variability to be explored. In other cases (e.g., grain-size sorting of a detrital component),
463 sample-to-sample variability is physically linked to a single EM with continuously variable
464 properties, rather than to a mechanical mixture of EMs with fixed properties. In these cases, FORC-
465 PCA generates two or more ‘fictive’ EMs that recreate inter-sample variations, but do not
466 correspond to fixed physical components of the system. Nevertheless, the mixing proportions of
467 ‘fictive’ EMs provide a useful co-ordinate system with which to quantify the extent of inter-sample
468 variation, and may be used to identify variation trends and clusters of behavior.

469

470 The ability to unmix up to four¹ EMs, each with their own distinct domain state, coercivity and
471 interaction field distribution, takes us beyond the routine characterization that is commonly
472 considered ‘good enough’ for most paleomagnetic studies, and largely addresses the ambiguities
473 (Roberts et al., 2018) involved in interpreting the widely used Day diagram (Day et al., 1977). The
474 new algorithm provides a full set of FORCs for each EM, which allows additional hysteresis
475 properties to be derived for each EM. In some cases, this additional information can be used to
476 check for consistency with the interpreted physical origin of each EM (e.g., if the FORC diagram of
477 the EM suggests non-interacting, uniaxial SD behavior, then M_r/M_s values close to 0.5 and B_{cr}/B_c
478 values close to 1 would be expected). Compared to a Day diagram, hysteresis ratios of extracted
479 EMs acquire enhanced physical meaning because the effects of mixing have been deconvolved. The
480 use of feasibility metrics reduces (but does not eliminate) the ambiguity involved in defining EMs
481 when the unmixing space is sampled incompletely. This development should help to make the
482 FORC-PCA method accessible to a wider audience. However, it should always be borne in mind
483 that feasibility metrics are only a guideline – good choices, as ever, rely on the expertise and
484 judgement of the user.

485

486 **9. Conclusions**

¹ There is no limit on the number of EMs that can be mathematically defined. However, beyond four EMs, visualization and interactive exploration of the unmixing space becomes impractical.

- 487 1. Our improved FORC-PCA algorithm addresses many of the outstanding issues with the initial
488 method of Lascu et al. (2015), including solving the linear mixing problem and providing the
489 ability to characterize SP EMs that are dominated by reversible magnetizations.
- 490 2. The new method enables both the reconstructed FORC magnetization surface and the
491 corresponding FORC diagram of each EM to be identified.
- 492 3. Access to the reconstructed FORC magnetization surface enables objective criteria to be defined
493 that guide the choice of physically realistic EMs. A mixture of robust criteria (e.g., saturation,
494 monotonicity, and crossing) and more flexible criteria (e.g., positivity) can be used to help
495 reduce the subjectivity of defining the unmixing space.
- 496 4. The method has been applied successfully to quantify synthetic binary mixtures with EMs with
497 contrasting domain states, and to aid interpretation of diagenetic trends in greigite-bearing
498 sedimentary environments.
- 499 5. The improved FORC-PCA algorithm provides a powerful method to discriminate between
500 populations of grains with different domain state, coercivity distribution, anisotropy type, and
501 interaction field distribution. The increased value of the information that this analysis yields far
502 outweighs the additional measurement time that is needed, providing a way to take routine rock
503 magnetic characterization far beyond the ambiguities of the widely used Day diagram.

504

505 **Acknowledgements**

506 We thank Andrew Newell and Ramon Egli for their helpful and constructive reviews of the
507 manuscript. This work was supported financially by the Australian Research Council through grant
508 DP160100805 and by the European Research Council under the European Union's Seventh
509 Framework Programme (FP/2007–2013)/ERC grant agreement number 320750. Samples of SD-V
510 and SD-MV mixtures were kindly provided by Bruce Moskowitz of the Institute for Rock
511 Magnetism, University of Minnesota. All data used in this study is available for download as part of
512 the online supporting information.

513

514 **References**

- 515 Carter-Stiglitz, B., Moskowitz, B., & Jackson, M. (2001). Unmixing magnetic assemblages and the
516 magnetic behavior of bimodal mixtures. *Journal of Geophysical Research: Solid Earth*,
517 *106*(B11), 26397. <https://doi.org/10.1029/2001JB000417>
- 518 Carvalho, C., & Muxworthy, A. (2006). Low-temperature first-order reversal curve (FORC)
519 diagrams for synthetic and natural samples. *Geochemistry, Geophysics, Geosystems*, *7*(9),
520 Q09003. <https://doi.org/10.1029/2006GC001299>
- 521 Carvalho, C., Dunlop, D., & Ozdemir, O. (2005). Experimental comparison of FORC and remanent
522 Preisach diagrams. *Geophysical Journal International*, *162*(3), 747–754. Retrieved from
523 <papers://2973f1d5-d696-469c-80fa-0e7ffe9241d4/Paper/p421>
- 524 Channell, J. E. T., Harrison, R. J., Lascu, I., McCave, I. N., Hibbert, F. D., & Austin, W. E. N.
525 (2016). Magnetic record of deglaciation using FORC-PCA, sortable-silt grain size, and
526 magnetic excursion at 26 ka, from the Rockall Trough (NE Atlantic). *Geochemistry*,
527 *Geophysics, Geosystems*, *17*(5), 1823–1841. <https://doi.org/10.1002/2016GC006300>
- 528 Day, R., Fuller, M., & Schmidt, V. A. (1977). Hysteresis properties of titanomagnetites: grain-size
529 and compositional dependence. *Physics of the Earth and Planetary Interiors*, *13*, 260–267.
- 530 Donnelly, C., Guizar-Sicairos, M., Scagnoli, V., Gliga, S., Holler, M., Raabe, J., & Heyderman, L.
531 J. (2017). Three-dimensional magnetization structures revealed with X-ray vector
532 nanotomography. *Nature*, *547*(7663), 328–331. <https://doi.org/10.1038/nature23006>
- 533 Dunlop, D. J. (2002a). Theory and application of the Day plot (Mrs/Ms versus Hcr/Hc) 1.
534 Theoretical curves and tests using titanomagnetite data. *Journal of Geophysical Research:*
535 *Solid Earth*, *107*(B3), 2056. <https://doi.org/10.1029/2001JB000486>
- 536 Dunlop, D. J. (2002b). Theory and application of the Day plot (Mrs/Ms versus Hcr/Hc) 2.
537 Application to data for rocks, sediments, and soils. *Journal of Geophysical Research: Solid*
538 *Earth*, *107*(B3), 2057. <https://doi.org/10.1029/2001JB000487>

- 539 Dunlop, D. J., & Carter-Stiglitz, B. (2006). Day plots of mixtures of superparamagnetic, single-
540 domain, pseudosingle-domain, and multidomain magnetites. *Journal of Geophysical Research:*
541 *Solid Earth*, 111(B12), B12S09. <https://doi.org/10.1029/2006JB004499>
- 542 Ebert, Y., Shaar, R., Emmanuel, S., Nowaczyk, N., & Stein, M. (2018). Overwriting of sedimentary
543 magnetism by bacterially mediated mineral alteration. *Geology*, 46(4), 2–5.
544 <https://doi.org/https://doi.org/10.1130/G39706.1>
- 545 Egli, R. (2004a). Characterization of Individual Rock Magnetic Components by Analysis of
546 Remanence Curves, 1. Unmixing Natural Sediments. *Studia Geophysica et Geodaetica*, 48(2),
547 391–446. <https://doi.org/10.1023/B:SGEG.0000020839.45304.6d>
- 548 Egli, R. (2004b). Characterization of individual rock magnetic components by analysis of
549 remanence curves. *Physics and Chemistry of the Earth*, 29(13–14), 851–867.
550 <https://doi.org/10.1016/j.pce.2004.04.001>
- 551 Egli, R. (2004c). Characterization of individual rock magnetic components by analysis of
552 remanence curves. 3. Bacterial magnetite and natural processes in lakes. *Physics and*
553 *Chemistry of the Earth*, 29(13–14), 869–884. <https://doi.org/10.1016/j.pce.2004.03.010>
- 554 Egli, R. (2006). Theoretical aspects of dipolar interactions and their appearance in first-order
555 reversal curves of thermally activated single-domain particles. *Journal of Geophysical*
556 *Research: Solid Earth*, 111(B12), B12S17. <https://doi.org/10.1029/2006JB004567>
- 557 Egli, R. (2013). VARIFORC: An optimized protocol for calculating non-regular first-order reversal
558 curve (FORC) diagrams. *Global and Planetary Change*, 110, 302–320.
559 <https://doi.org/10.1016/j.gloplacha.2013.08.003>
- 560 Egli, R., Chen, A. P., Winklhofer, M., Kodama, K. P., & Horng, C.-S. (2010). Detection of
561 noninteracting single domain particles using first-order reversal curve diagrams. *Geochemistry,*
562 *Geophysics, Geosystems*, 11(1), Q01Z11. <https://doi.org/10.1029/2009GC002916>
- 563 Egli, R., & Winklhofer, M. (2014). Recent developments on processing and interpretation aspects
564 of first-order reversal curves (FORC). *Proc. Kazan. Univ.*, 156(January), 14–53.

- 565 Evans, M. E., & Heller, F. (2003). *Environmental magnetism principles and applications of*
566 *enviromagnetics*. Academic Press, Amsterdam.
- 567 Florindo, F., Karner, D., Marra, F., Renne, P. R., Roberts, A. P., & Weaver, R. (2007).
568 Radioisotopic age constraints for Glacial Terminations IX and VII from aggradational sections
569 of the Tiber River delta in Rome, Italy. *Earth and Planetary Science Letters*, 256(1–2), 61–80.
570 <https://doi.org/10.1016/j.epsl.2007.01.014>
- 571 Franke, C., Frederichs, T., & Dekkers, M. J. (2007). Efficiency of heavy liquid separation to
572 concentrate magnetic particles. *Geophysical Journal International*, 170(3), 1053–1066.
573 <https://doi.org/10.1111/j.1365-246X.2007.03489.x>
- 574 Harrison, R. J., & Feinberg, J. M. (2008). FORCinel: An improved algorithm for calculating first-
575 order reversal curve distributions using locally weighted regression smoothing. *Geochemistry,*
576 *Geophysics, Geosystems*, 9(5), Q05016. <https://doi.org/10.1029/2008GC001987>
- 577 Harrison, R. J., & Lascu, I. (2014). FORCulator: A micromagnetic tool for simulating first-order
578 reversal curve diagrams. *Geochemistry, Geophysics, Geosystems*, 15(12), 4671–4691.
579 <https://doi.org/10.1002/2014GC005582>
- 580 Heslop, D. (2015). Numerical strategies for magnetic mineral unmixing. *Earth-Science Reviews*,
581 150, 256–284. <https://doi.org/10.1016/j.earscirev.2015.07.007>
- 582 Heslop, D., Roberts, A. P., & Chang, L. (2014). Characterizing magnetofossils from first-order
583 reversal curve (FORC) central ridge signatures. *Geochemistry, Geophysics, Geosystems*, 15(6),
584 2170–2179. <https://doi.org/10.1002/2014GC005291>
- 585 Heslop, D., & Dillon, M. (2007). Unmixing magnetic remanence curves without a priori
586 knowledge. *Geophysical Journal International*, 170(2), 556–566.
587 <https://doi.org/10.1111/j.1365-246X.2007.03432.x>
- 588 Heslop, D., & Roberts, A. P. (2012a). Estimation of significance levels and confidence intervals for
589 first-order reversal curve distributions. *Geochemistry Geophysics Geosystems*, 13(12),
590 Q12Z40. <https://doi.org/10.1029/2012GC004115>

- 591 Heslop, D., & Roberts, A. P. (2012b). A method for unmixing magnetic hysteresis loops. *Journal of*
592 *Geophysical Research: Solid Earth*, 117(B3), B03103. <https://doi.org/10.1029/2011JB008859>
- 593 Jolliffe, I. T. (2002) Principal Component Analysis. Springer-Verlag, New York, pp. 488.
- 594 Kruiver, P. P., Dekkers, M. J., & Heslop, D. (2001). Quantification of magnetic coercivity
595 components by the analysis of acquisition curves of isothermal remanent magnetisation. *Earth*
596 *and Planetary Science Letters*, 189(3–4), 269–276. [https://doi.org/10.1016/S0012-](https://doi.org/10.1016/S0012-821X(01)00367-3)
597 821X(01)00367-3
- 598 Lagroix, F., & Guyodo, Y. (2017). A new tool for separating the magnetic mineralogy of complex
599 mineral assemblages from low temperature magnetic behavior. *Frontiers in Earth Science*, 5,
600 1–11. <https://doi.org/10.3389/feart.2017.00061>
- 601 Lascu, I., Banerjee, S. K., & Berquó, T. S. (2010). Quantifying the concentration of ferrimagnetic
602 particles in sediments using rock magnetic methods. *Geochemistry, Geophysics, Geosystems*,
603 11(8), Q08Z19. <https://doi.org/10.1029/2010GC003182>
- 604 Lascu, I., Harrison, R. J., Li, Y., Muraszko, J. R., Channell, J. E. T., Piotrowski, A. M., & Hodell,
605 D. A. (2015). Magnetic unmixing of first-order reversal curve diagrams using principal
606 component analysis. *Geochemistry, Geophysics, Geosystems*, 16(9), 2900–2915.
607 <https://doi.org/10.1002/2015GC005909>
- 608 Li, J., Wu, W., Liu, Q., & Pan, Y. (2012). Magnetic anisotropy, magnetostatic interactions and
609 identification of magnetofossils. *Geochemistry, Geophysics, Geosystems*, 13(12), Q10Z51.
610 <https://doi.org/10.1029/2012GC004384>
- 611 Lindquist, A. K., Feinberg, J. M., Harrison, R. J., Loudon, J. C., & Newell, A. J. (2015). Domain
612 wall pinning and dislocations: Investigating magnetite deformed under conditions analogous to
613 nature using transmission electron microscopy. *Journal of Geophysical Research: Solid Earth*,
614 120(3), 1415–1430. <https://doi.org/10.1002/2014JB011335>
- 615 Liu, Q., Roberts, A. P., Larrasoña, J. C., Banerjee, S. K., Guyodo, Y., Tauxe, L., & Oldfield, F.
616 (2012). Environmental magnetism: Principles and applications. *Reviews of Geophysics*, 50(4),

- 617 RG4002. <https://doi.org/10.1029/2012RG000393>
- 618 Ludwig, P., Egli, R., Bishop, S., Chernenko, V., Frederichs, T., Rugel, G., ... Orgeira, M. J. (2013).
619 Characterization of primary and secondary magnetite in marine sediment by combining
620 chemical and magnetic unmixing techniques. *Global and Planetary Change*, *110*, 321–339.
621 <https://doi.org/10.1016/j.gloplacha.2013.08.018>
- 622 Maxbauer, D. P., Feinberg, J. M., Fox, D. L., & Nater, E. A. (2017). Response of pedogenic
623 magnetite to changing vegetation in soils developed under uniform climate, topography, and
624 parent material. *Scientific Reports*, *7*(1), 17575. <https://doi.org/10.1038/s41598-017-17722-2>
- 625 Moskowitz, B., Frankel, R. B., & Bazylinski, D. A. (1993). Rock magnetic criteria for the detection
626 of biogenic magnetite. *Earth and Planetary Science Letters*, *120*, 283–300.
- 627 Muxworthy, A. R., King, J., & Helsop, D. (2005). Assessing the ability of first-order reversal curve
628 (FORC) diagrams to unravel complex magnetic signals. *Journal of Geophysical Research*,
629 *110*(B1), B01105. <https://doi.org/10.1029/2004JB003195>
- 630 Muxworthy, A., Heslop, D., & Williams, W. (2004). Influence of magnetostatic interactions on
631 first-order-reversal-curve (FORC) diagrams: a micromagnetic approach. *Geophysical Journal
632 International*, *158*(3), 888–897. <https://doi.org/10.1111/j.1365-246X.2004.02358.x>
- 633 Newell, A. J. (2005). A high-precision model of first-order reversal curve (FORC) functions for
634 single-domain ferromagnets with uniaxial anisotropy. *Geochemistry, Geophysics, Geosystems*,
635 *6*(5), Q05010. <https://doi.org/10.1029/2004GC000877>
- 636 Pike, C. R. (2003). First-order reversal-curve diagrams and reversible magnetization. *Physical
637 Review B*, *68*(10), 104424. <https://doi.org/10.1103/PhysRevB.68.104424>
- 638 Pike, C. R., Roberts, A. P., & Verosub, K. L. (1999). Characterizing interactions in fine magnetic
639 particle systems using first order reversal curves. *Journal of Applied Physics*, *85*(9), 6660–
640 6667.
- 641 Pike, C., Roberts, A., Dekkers, M., & Verosub, K. (2001a). An investigation of multi-domain
642 hysteresis mechanisms using FORC diagrams. *Physics of the Earth and Planetary Interiors*,

- 643 126, 11–25.
- 644 Pike, C. R., Roberts, A. P., & Verosub, K. L. (2001b). First-order reversal curve diagrams and
645 thermal relaxation effects in magnetic particles. *Geophysical Journal International*, 145, 721–
646 730.
- 647 Plaza-Morlote, M., Rey, D., Santos, J. F., Ribeiro, S., Heslop, D., Bernabeu, A., ... Martíns, V.
648 (2017). Southernmost evidence of large European Ice Sheet-derived freshwater discharges
649 during the Heinrich Stadials of the Last Glacial Period (Galician Interior Basin, Northwest
650 Iberian Continental Margin). *Earth and Planetary Science Letters*, 457, 213–226.
651 <https://doi.org/10.1016/j.epsl.2016.10.020>
- 652 Roberts, A., Chang, L., Rowan, C., Horng, C., & Florindo, F. (2011). Magnetic properties of
653 sedimentary greigite (Fe₃S₄): An update. *Reviews of Geophysics*, 49(1), RG1002.
- 654 Roberts, A. P., Tauxe, L., Heslop, D., Zhao, X., & Jiang, Z. (2018). A Critical Appraisal of the
655 “Day” Diagram. *Journal of Geophysical Research: Solid Earth*.
656 <https://doi.org/10.1002/2017JB015247>
- 657 Roberts, A. P., Heslop, D., Zhao, X., & Pike, C. R. (2014). Understanding fine magnetic particle
658 systems through use of first-order reversal curve diagrams. *Reviews of Geophysics*, 52(4), 557–
659 602. <https://doi.org/10.1002/2014RG000462>
- 660 Roberts, A. P., Pike, C. R., & Verosub, K. L. (2000). First order reversal curve diagrams: A new
661 tool for characterising the magnetic properties of natural samples. *Journal of Geophysical
662 Research: Solid Earth*, 105(B12), 28461–28475.
- 663 Roberts, A. P., Almeida, T. P., Church, N. S., Harrison, R. J., Heslop, D., Li, Y., ... Zhao, X.
664 (2017). Resolving the Origin of Pseudo-Single Domain Magnetic Behavior. *Journal of
665 Geophysical Research: Solid Earth*, 122(12), 9534–9558.
666 <https://doi.org/10.1002/2017JB014860>
- 667 Robertson, D. J., & France, D. E. (1994). Discrimination of remanence-carrying minerals in
668 mixtures, using isothermal remanent magnetisation acquisition curves. *Physics of the Earth*

- 669 *and Planetary Interiors*, 82(3–4), 223–234. [https://doi.org/10.1016/0031-9201\(94\)90074-4](https://doi.org/10.1016/0031-9201(94)90074-4)
- 670 Rowan, C. J., & Roberts, A. P. (2006). Magnetite dissolution, diachronous greigite formation, and
 671 secondary magnetizations from pyrite oxidation: Unravelling complex magnetizations in
 672 Neogene marine sediments from New Zealand. *Earth and Planetary Science Letters*, 241,
 673 119–137. <https://doi.org/10.1016/j.epsl.2005.10.017>
- 674 Zhao, X., Roberts, A. P., Heslop, D., Paterson, G. A., Li, Y., & Li, J. (2017). Magnetic domain state
 675 diagnosis using hysteresis reversal curves. *Journal of Geophysical Research: Solid Earth*,
 676 122(7), 4767–4789. <https://doi.org/10.1002/2016JB013683>
- 677 Zhao, X., Heslop, D., & Roberts, A. P. (2015). A protocol for variable-resolution first-order reversal
 678 curve measurements. *Geochemistry, Geophysics, Geosystems*, 16(5), 1364–1377.
 679 <https://doi.org/10.1002/2014GC005680>

680

681 **Figure Captions**

682

683 **Figure 1.** Experimental FORC diagrams for SD-MD mixtures. Smoothing performed using FORCinel with
 684 VARIFORC parameters $S_{c0} = 7$, $S_{c1} = 7$, $S_{b0} = 5$, $S_{b1} = 7$, $\lambda_c = 0.1$, and $\lambda_b = 0.1$. Mixing proportions are (a)
 685 100%, MD 0% SD, (b) 88%, MD 12% SD, and (c) 8% MD, 92% SD.

686

687 **Figure 2.** Experimental FORC diagrams for SD-V mixtures. Smoothing performed using FORCinel with
 688 VARIFORC parameters $S_{c0} = 7$, $S_{c1} = 12$, $S_{b0} = 5$, $S_{b1} = 12$, $\lambda_c = 0$, and $\lambda_b = 0$. Mixing proportions are (a)
 689 100% V, 0% SD, (b) 80% V, 20% SD, and (c) 5% V, 95% SD.

690

691 **Figure 3.** Experimental FORC diagrams for V-MD mixtures. Smoothing performed using FORCinel with
 692 VARIFORC parameters $S_{c0} = 7$, $S_{c1} = 7$, $S_{b0} = 5$, $S_{b1} = 7$, $\lambda_c = 0.1$, and $\lambda_b = 0.1$. Mixing proportions are (a)
 693 100% V, 0% MD, (b) 49% V, 51% MD, and (c) 0% V, 100% MD.

694

695 **Figure 4.** Comparison of mixing proportions derived from FORC-PCA (EM) using the method of Lascu et
 696 al. (2015) with known mass proportions of end members in synthetic mixtures (EM*) of (a) SD-MD, (b) SD-

697 V, and (c) V-MD magnetite particles, respectively. Solid lines are fits to the data using eqn. 6, which yield f
 698 = 2.72, 1.52, and 2.69, respectively.

699

700 **Figure 5.** FORC-PCA analysis of V-MD mixtures using the new unmixing algorithm. **(a-b)** Reconstructed
 701 FORC diagrams for EM1 (V) and EM2 (MD). **(c-d)** Reconstructed FORCs for EM1 and EM2. **(e)** PC score
 702 plot for a binary unmixing space between EM1 and EM2 (indicated by arrows). Diamonds illustrate the
 703 scores of individual samples. **(f)** Comparison of mixing proportions extracted using the new algorithm with
 704 the known mass proportions of end members in the synthetic mixture. The solid line indicates a one-to-one
 705 relationship.

706

707 **Figure 6.** FORC-PCA analysis of MD-SD mixtures using the new algorithm. **(a-c)** Reconstructed FORC
 708 diagrams for EM1 (MD), EM2 (SD high coercivity), and SD (low coercivity), respectively. **(d-f)**
 709 Reconstructed FORCs for EM1, EM2, and EM3, respectively. **(g)** PC score plot for a ternary unmixing space
 710 between EM1, EM2, and EM3 (black triangle). Diamonds illustrate the scores of individual samples.
 711 Contour lines represent the combined feasibility metric, m , for the saturation, monotonicity, and crossing
 712 metrics. **(h)** Comparison of mixing proportions extracted using the new algorithm with the known mass
 713 proportions of end members in the synthetic mixture. The solid line represents a one-to-one relationship.

714

715 **Figure 7.** FORC-PCA analysis of V-SD mixtures using the new algorithm. **(a-c)** Reconstructed FORC
 716 diagrams for EM1 (V), EM2 (SD high coercivity) and SD (low coercivity), respectively. **(d-f)** Reconstructed
 717 FORCs for EM1, EM2, and EM3, respectively. **(g)** PC score plot for a ternary unmixing space between
 718 EM1, EM2, and EM3 (black triangle). Diamonds illustrate the scores of individual samples. Contour lines
 719 represent the combined feasibility metric for the saturation, monotonicity, and crossing metrics. **(h)**
 720 Comparison of mixing proportions extracted using the new algorithm with the known mass proportions of
 721 end members in the synthetic mixture. The solid line represents a one-to-one relationship.

722

723 **Figure 8.** FORC-PCA analysis of greigite-bearing clay samples from the Tiber River, Rome (Florindo et al.
 724 (2007)). **(a-c)** Reconstructed FORC diagrams for EM1 (moderately interacting SD greigite), EM2 (strongly
 725 interacting SD greigite) and EM3 (SP/SD greigite), respectively. **(d-f)** Reconstructed FORCs for EM1, EM2,

726 and EM3. The arrow in (f) indicates the downward-inflected response at the start of each FORC. Although
727 this phenomenon is associated partially here with viscous SP behavior (Pike et al., 2001b), it is also likely to
728 be partially an instrumental artefact in this case. (g) PC score plot for a ternary unmixing space between
729 EM1, EM2, and EM3 (black triangle). Diamonds illustrate the scores of individual samples. Contour lines
730 represent the combined feasibility metric for the saturation, monotonicity, and crossing metrics. The inset is
731 an illustration of contours for the positivity metric. (h) Ternary diagram for the extracted proportions of
732 EM1, EM2, and EM3. The blue line illustrates the two dominant mixing trends (EM3-EM2 and EM2-EM1).
733 An example experimental FORC diagram for a mixture of strongly interacting and ~35% viscous SP/SD
734 greigite is indicated by the arrow.

Figure 1.

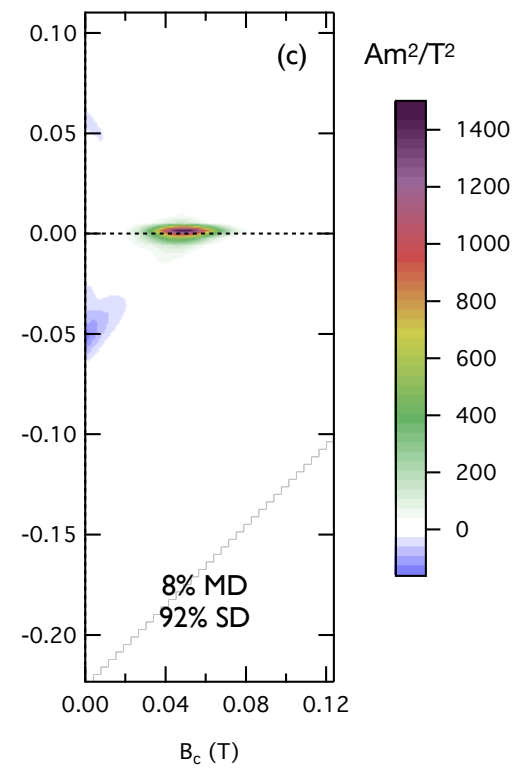
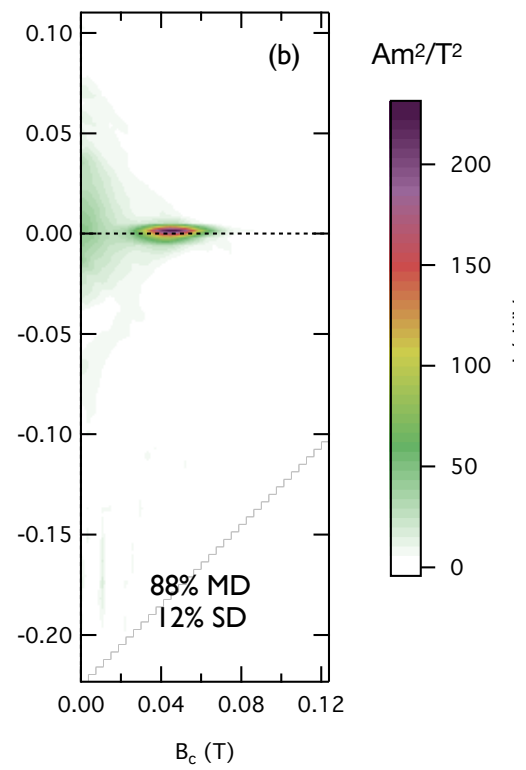
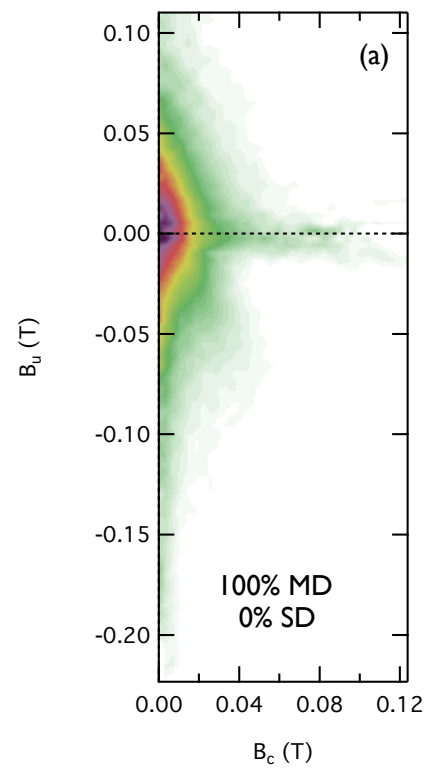


Figure 2.

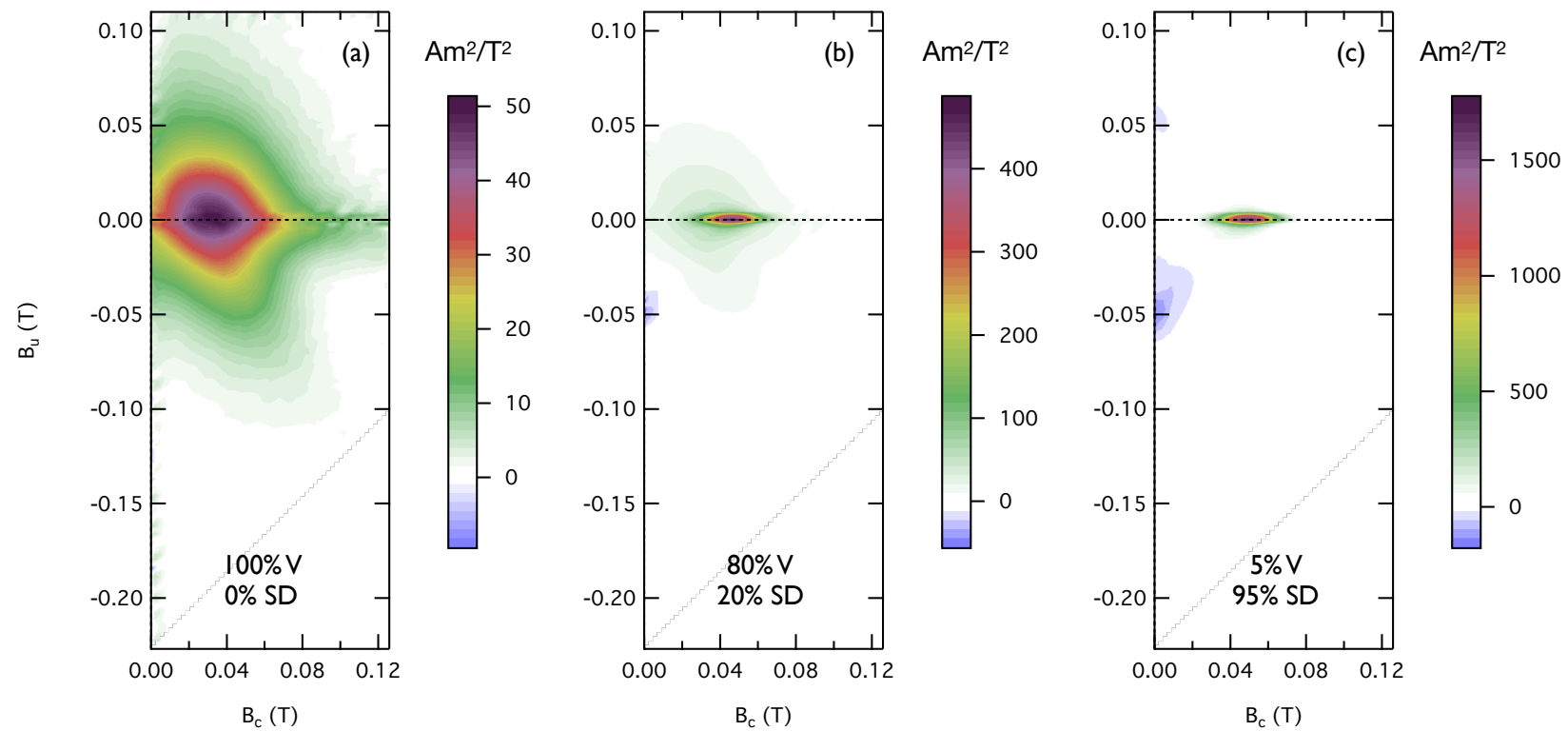


Figure 3.

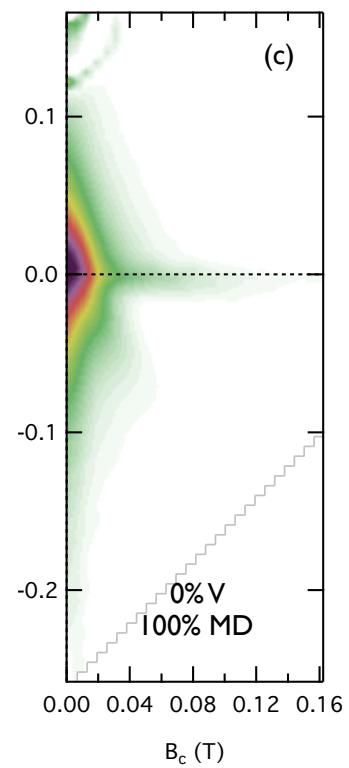
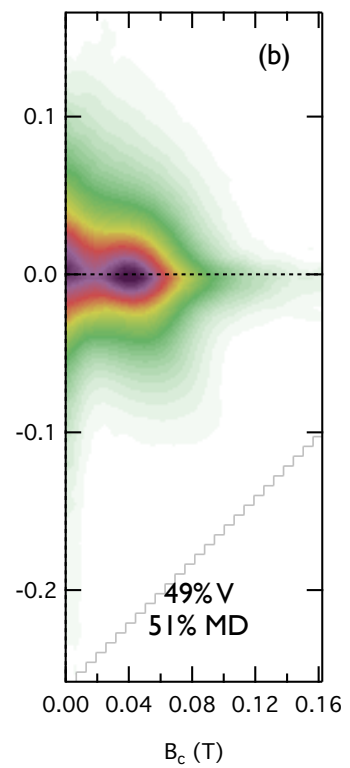
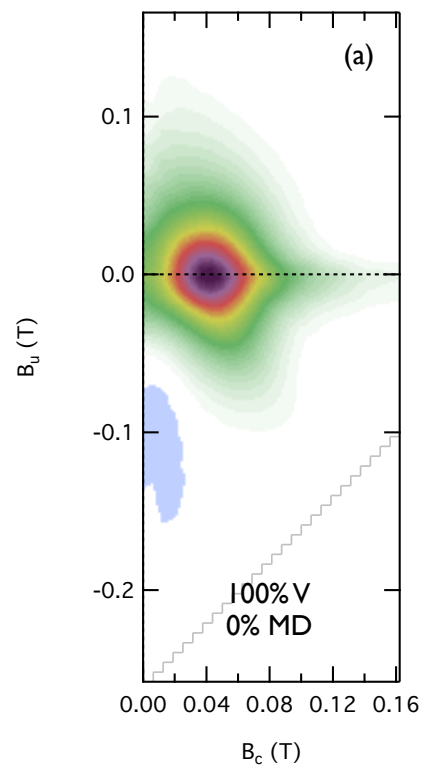


Figure 4.

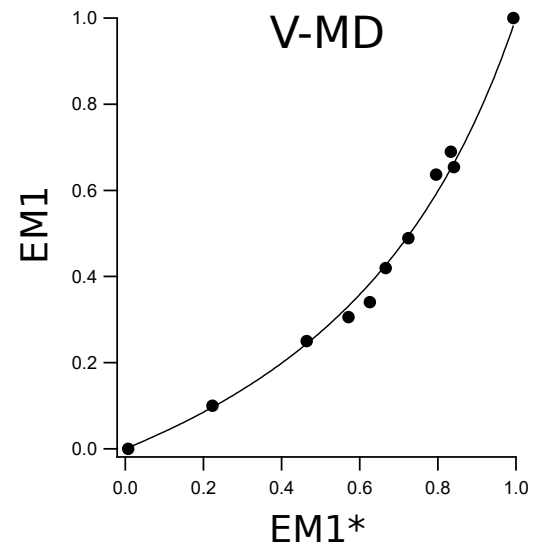
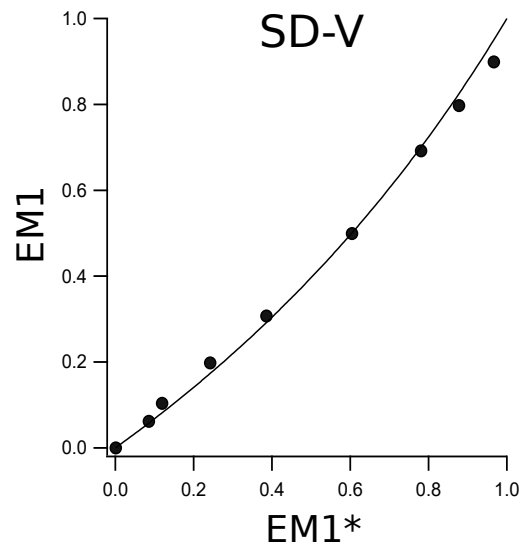
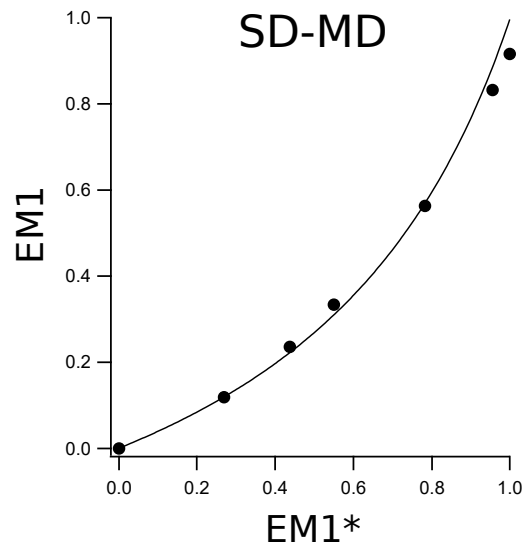


Figure 5.

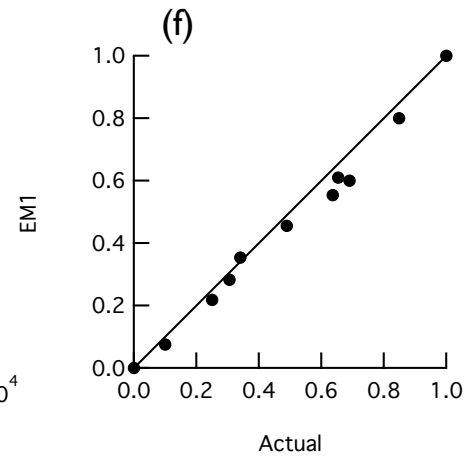
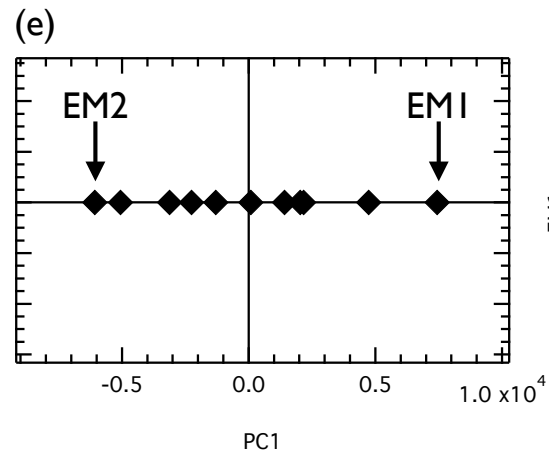
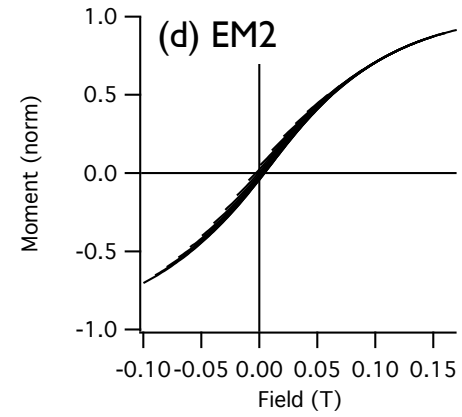
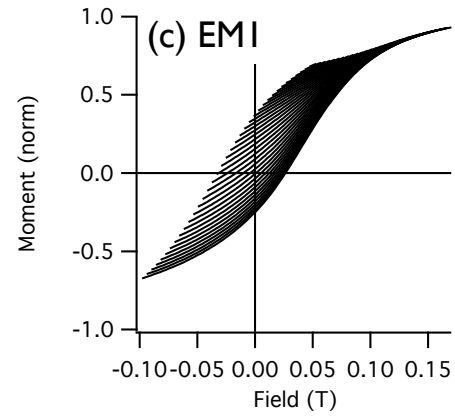
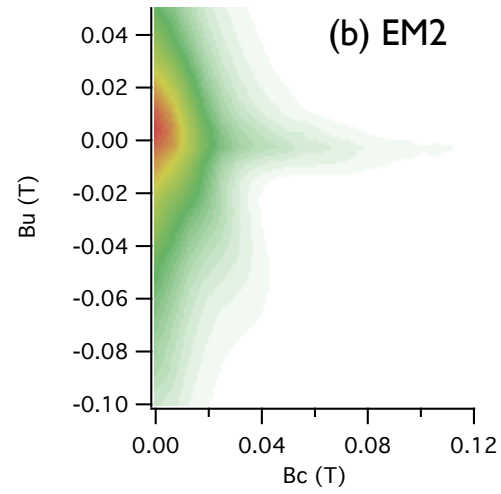
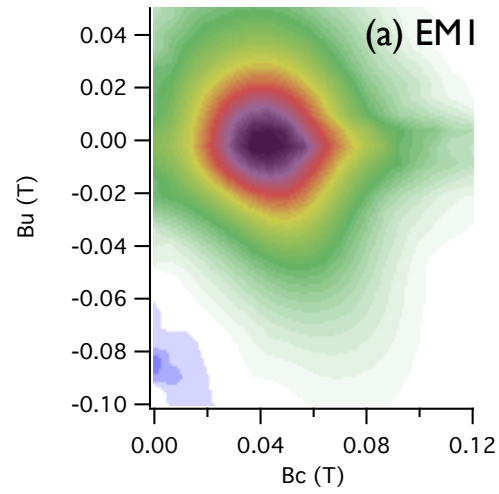


Figure 6.

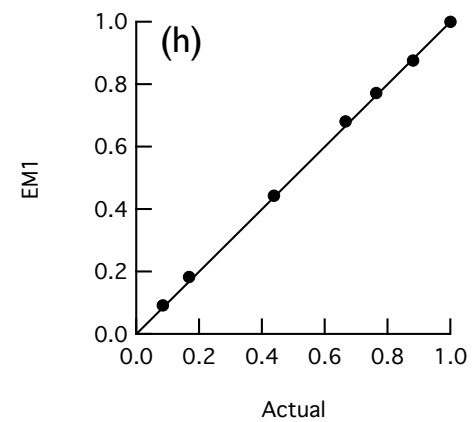
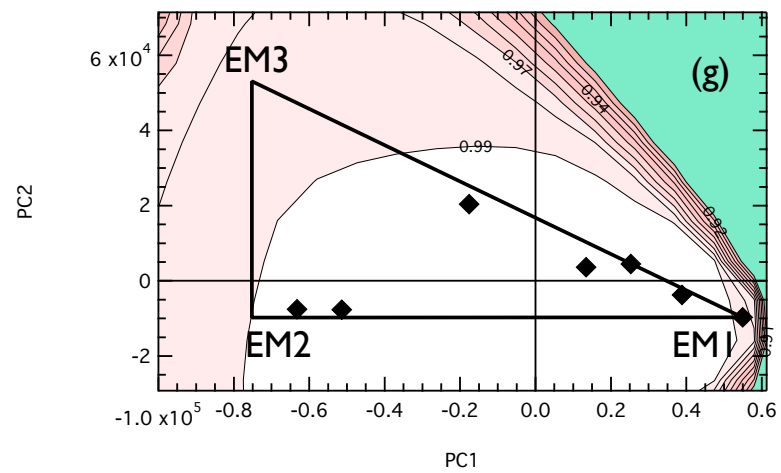
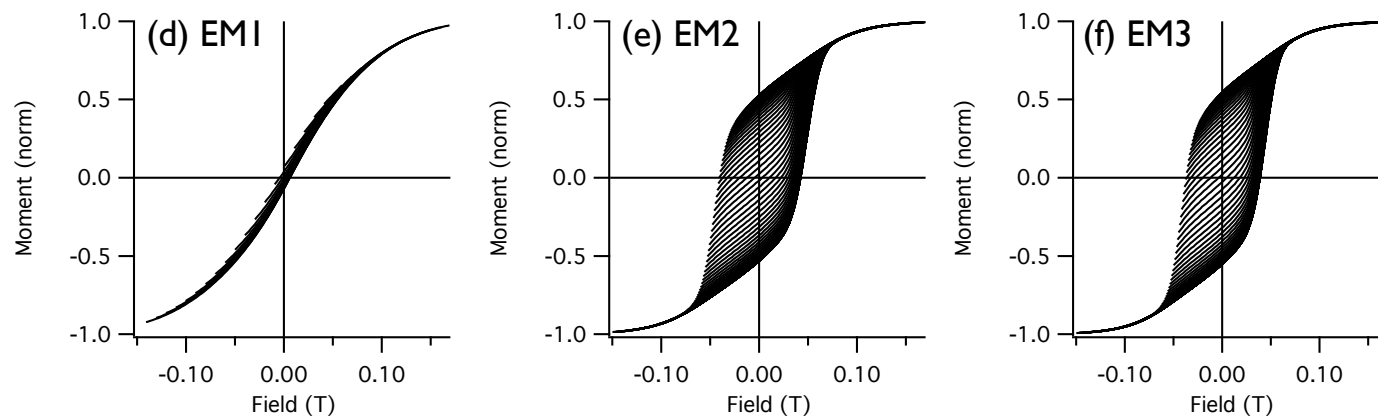
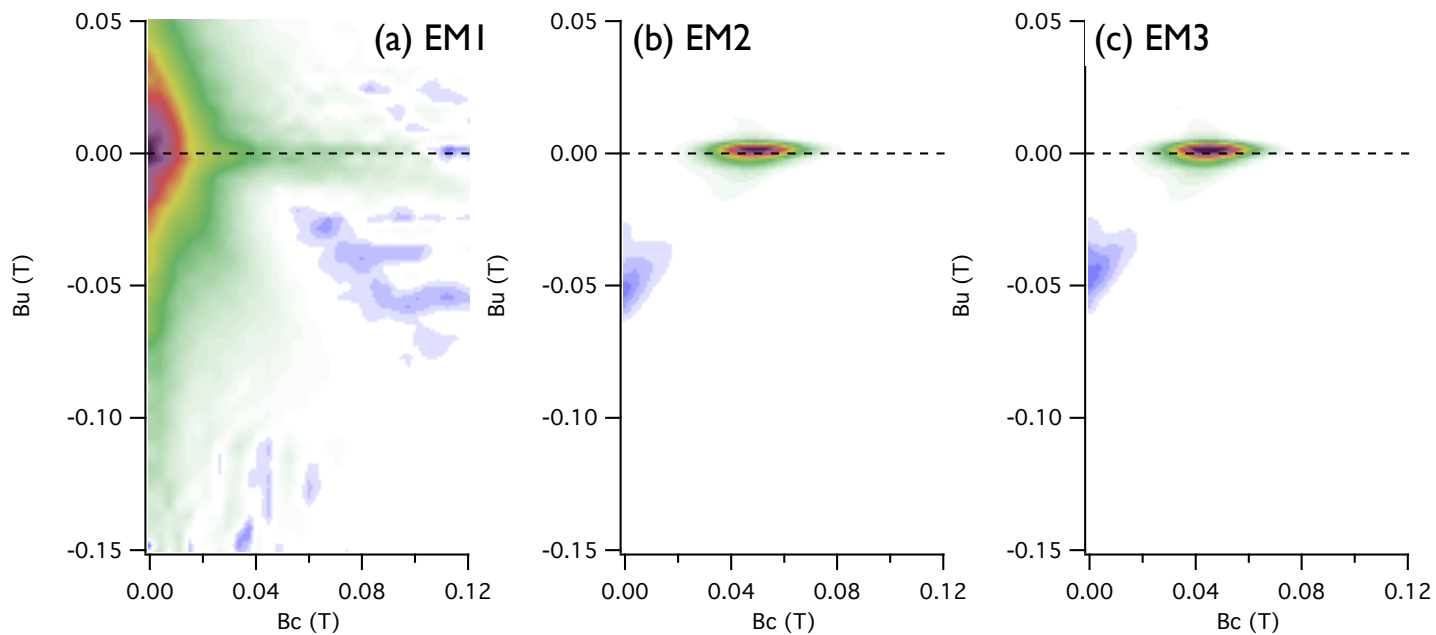


Figure 7.

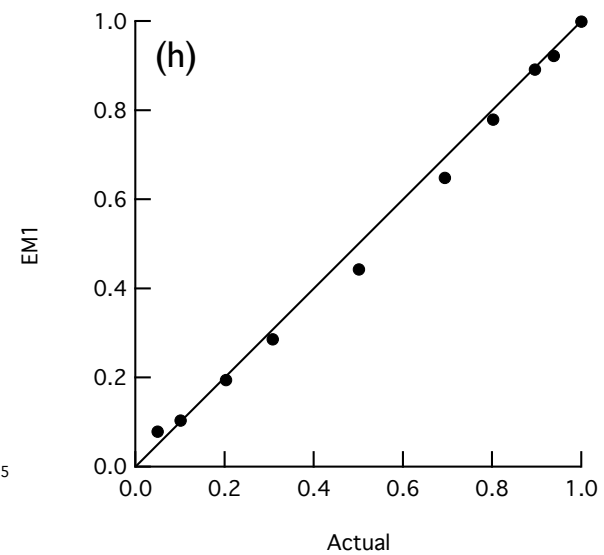
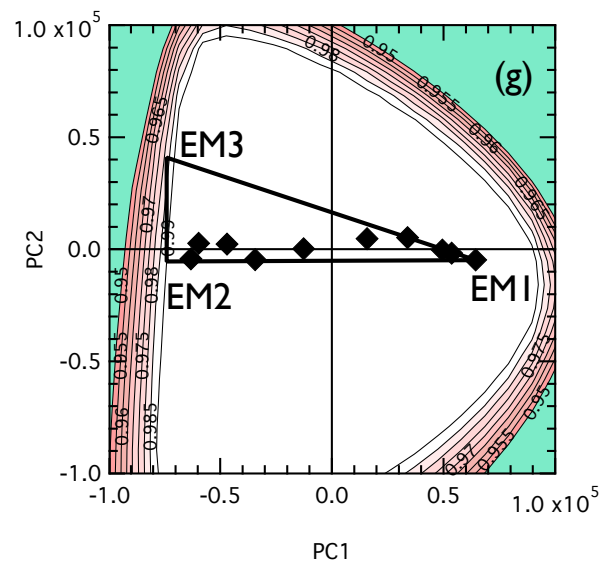
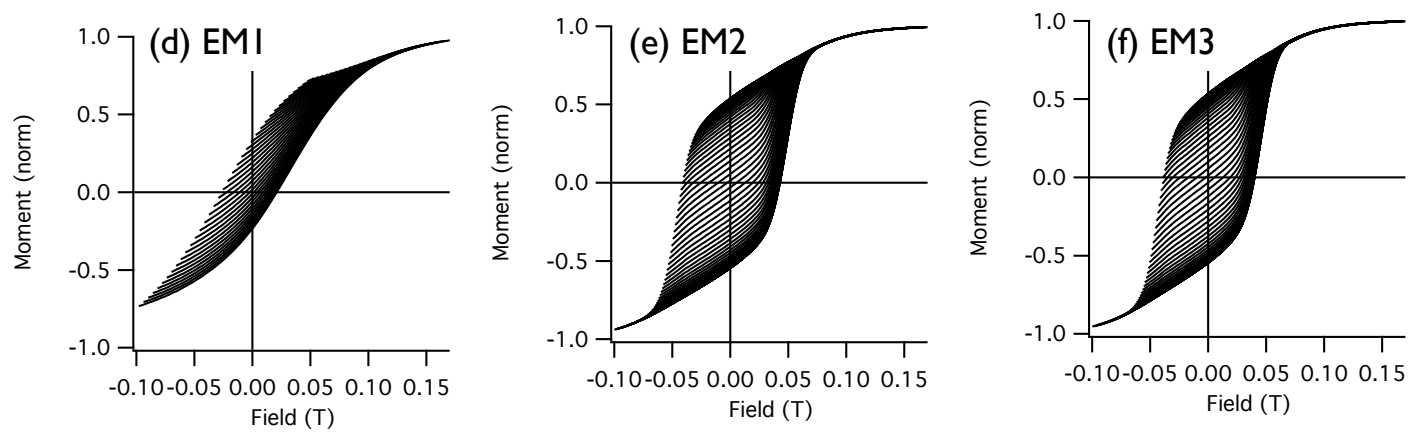
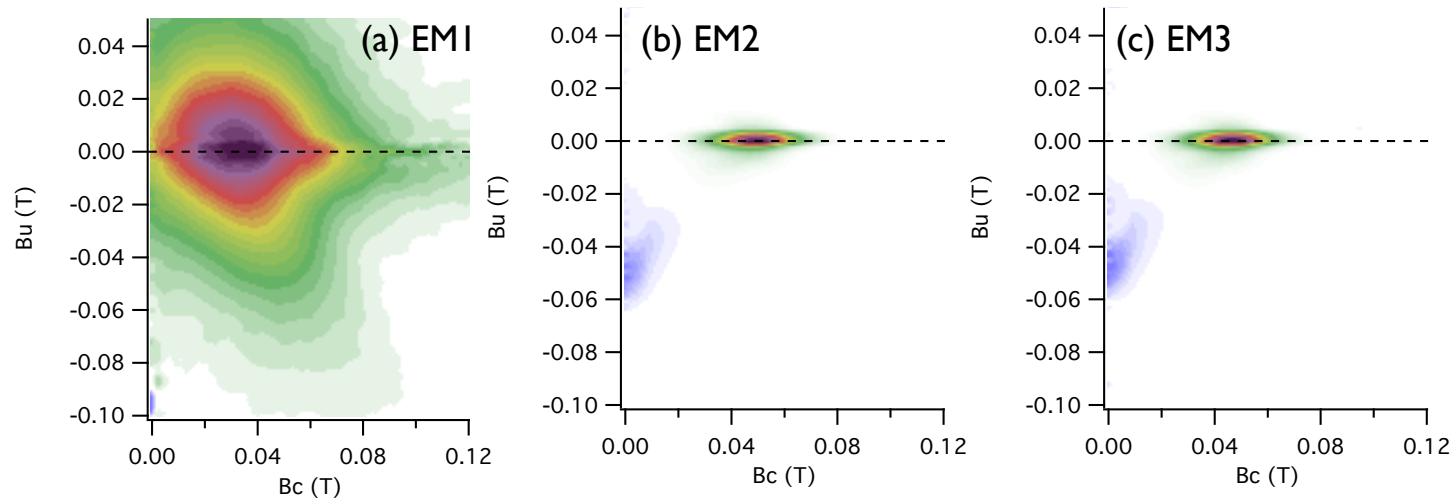


Figure 8.

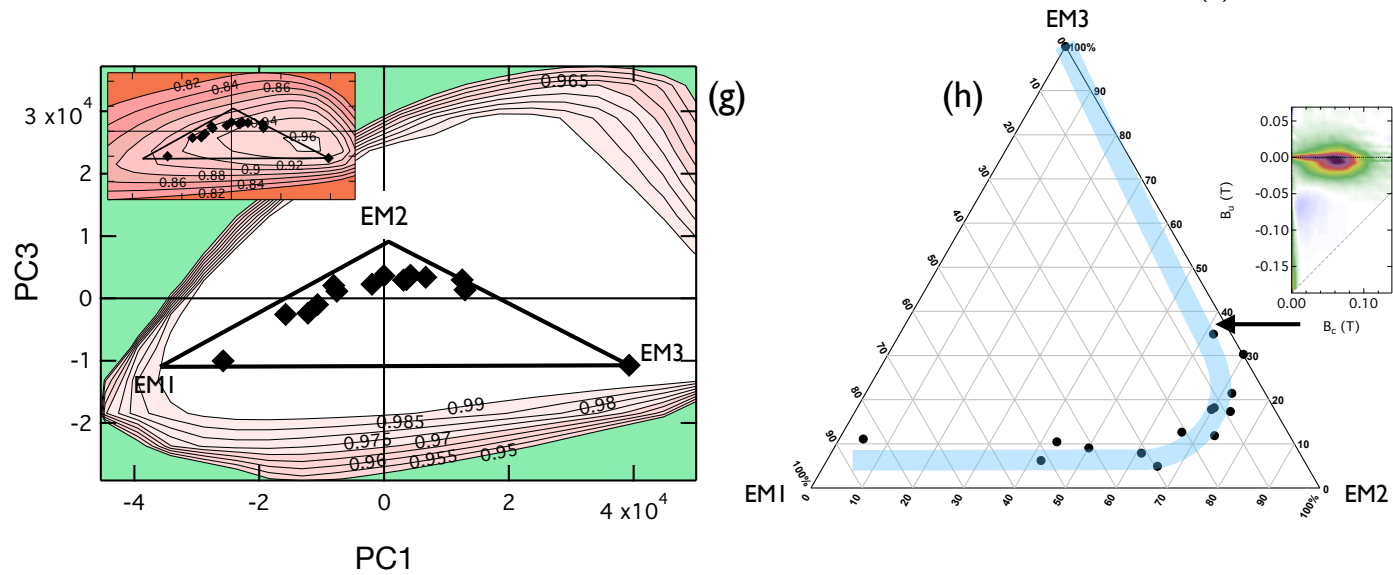
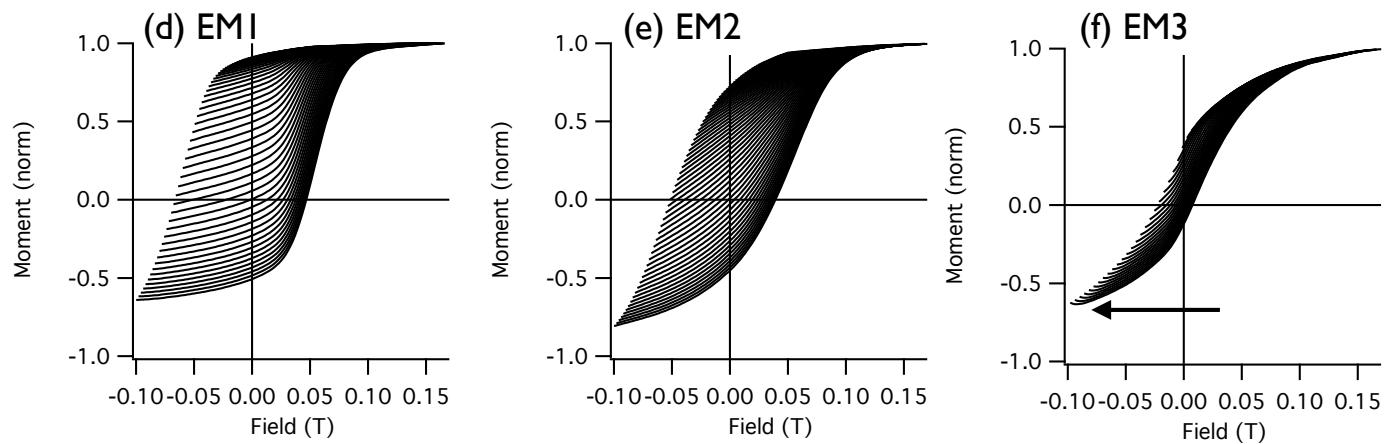
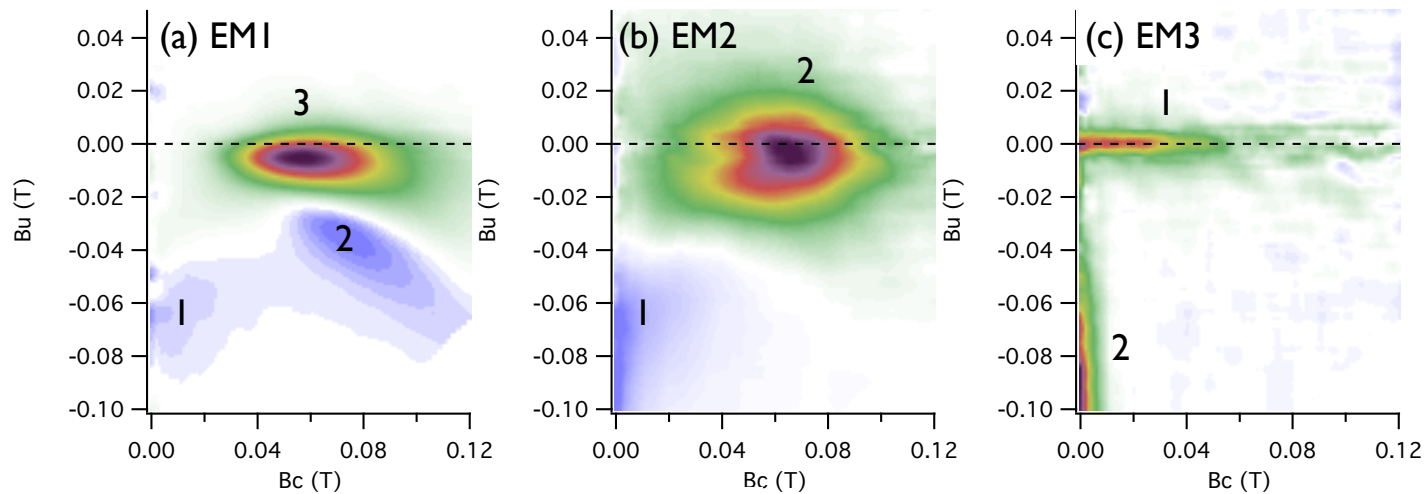


Table 1. Mass fractions of endmembers in measured samples			
Sample	SD	V	MD
V-MD			
wm_1		1.00	0.00
wm_2		0.85	0.15
wm_3		0.65	0.35
wm_4		0.64	0.36
wm_5		0.49	0.51
wm_6		0.69	0.31
wm_7		0.31	0.69
wm_8		0.34	0.66
wm_9		0.10	0.90
wm_10		0.00	1.00
wm_11		0.25	0.75
wm_12		0.42	0.58
SD-V			
w30_1	0.00	1.00	
w30_2	0.50	0.50	
w30_3	0.31	0.69	
w30_4	0.20	0.80	
w30_5	0.10	0.90	
w30_6	0.06	0.94	
w30_7	0.69	0.31	
w30_8	0.80	0.20	
w30_9	0.90	0.10	
w30_10	0.95	0.05	
SD-MD			
w14_1	0.00		1.00
w14_2	0.56		0.44
w14_3	0.33		0.67
w14_4	0.24		0.76
w14_5	0.12		0.88
w14_6	0.83		0.17
w14_7	0.92		0.08

Table 2. Summary of hysteresis properties for measured samples

Sample	H_c (mT)	H_{cr} (mT)	H_{cr}/H_c	M_r/M_s
V-MD				
wm_1	31.35	52.56	1.68	0.29
wm_2	28.61	52.23	1.83	0.26
wm_3	23.11	79.32	3.43	0.20
wm_4	21.05	49.43	2.35	0.19
wm_5	18.25	48.59	2.66	0.16
wm_6	22.89	48.59	2.12	0.20
wm_7	13.70	46.04	3.36	0.12
wm_8	15.18	46.82	3.08	0.13
wm_9	2.59	34.53	13.36	0.03
wm_10	4.39	23.70	5.40	0.04
wm_11	11.11	42.61	3.83	0.09
wm_12	16.35	47.93	2.93	0.14
SD-V				
w30_1	24.20	45.26	1.87	0.02
w30_2	41.17	52.96	1.29	0.35
w30_3	35.85	52.79	1.47	0.29
w30_4	31.75	52.11	1.64	0.26
w30_5	27.55	50.56	1.84	0.23
w30_6	26.63	50.01	1.88	0.23
w30_7	41.34	49.42	1.20	0.40
w30_8	41.62	48.71	1.17	0.43
w30_9	41.78	48.63	1.16	0.47
w30_10	41.88	47.61	1.14	0.48
SD-MD				
w14_1	5.47	26.11	4.77	0.04
w14_2	34.24	49.43	1.44	0.27
w14_3	21.26	48.75	2.29	0.16
w14_4	16.52	47.85	2.90	0.13
w14_5	10.57	44.81	4.24	0.09
w14_6	41.45	49.35	1.19	0.40
w14_7	42.80	49.49	1.16	0.45

Table 3. Summary of hysteresis properties for extracted EMs

V-MD				
	H_c (mT)	H_{cr}	H_{cr}/H_c	M_r/M_s
EM1	33.0	52.6	1.6	0.38
EM2	4.6	24.5	5.3	0.044
V (obs)	31.0	52.9	1.7	0.37
MD (obs)	4.5	24.4	5.5	0.045
SD-V				
EM1	27.0	49.0	1.8	0.34
EM2	42.0	48.0	1.1	0.54
EM3	39.0	45.0	1.2	0.54
SD (95%)	41.0	48.0	1.2	0.52
V (obs)	24	48.5	2.0	0.34
SD-MD				
EM1	6.0	26.8	4.5	0.07
EM2	41.0	47.5	1.2	0.53
EM3	37.5	43.5	1.2	0.55
SD (92%)	39.7	47.5	1.2	0.49
MD (obs)	5.2	26.9	5.1	0.07