

2 **Have natural lake expansion and landscape inundation**
3 **resulted in mercury increases in flooded lakes of the Great**
4 **Slave Lowlands (Northwest Territories, Canada)?**

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9 **Abstract** The inundation of terrestrial vegetation
10 following landscape flooding is an important potential
11 source of mercury to aquatic ecosystems, and may
12 modify mercury cycling, such as through increased
13 methylation. In the Great Slave Lowlands of Canada's
14 Northwest Territories, remarkable landscape flooding
15 has occurred over the recent past, which is the most
16 notable in at least the last several centuries. The
17 potential for this flooding to increase inorganic
18 mercury flux to the lakes of the region has not yet
19 been explored. In this study we used sediment cores
20 from five lakes experiencing a range of recently
21 documented lake expansion to test whether inundation
22 of terrestrial areas has increased the total mercury
23 concentrations in sediments, and resulted in increased
24 total mercury flux. Increases in sedimentary mercury

concentrations and fluxes in sediment cores from the 25
expanding lakes were relatively small and within the 26
range of non-expanded systems, suggesting that, to 27
date, flooding has not resulted in major total mercury 28
enrichment, unlike in experimental and natural reser- 29
voir impoundments. The potential for increased 30
methylation of existing inorganic mercury following 31
expansion was not explored in this paper because 32
methylmercury is dynamic in sediments and does not 33
preserve well, but is an important consideration for 34
future work. 35

Keywords Climate change · Contaminants · 36
Flooding · Lake sediments · Paleolimnology · Mercury 37

Introduction 38

Increased contaminant exposure of ecosystems 39
through anthropogenic inputs represents a major 40
stressor globally. Mercury (Hg), is a naturally occur- 41
ring element, released in large quantities by human 42
activities beyond the historical, pre-industrial range. 43
Because of the capacity for long-range transport of 44
mercury in the atmosphere, this contaminant is of 45
particular interest in remote northern regions, which 46
may lack direct sources of anthropogenic pollution. 47
Landscape changes have the potential to alter the 48
movement and mobility of mercury, with concomitant 49
impacts to ecosystems and foodwebs. As mercury 50

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51 stored in soils is primarily divalent Hg(II), changing
 52 redox conditions, and microbial activity in response to
 53 saturation of soils is an important control on total
 54 mercury mobility to surface waters, the methylation to
 55 bioaccumulating MeHg, and the formation of gaseous
 56 Hg⁰, and thus its potential flux to the atmosphere
 57 (Poulin et al. 2016). Hydrological changes resulting in
 58 the inundation of terrestrial terrain, for example during
 59 the creation of reservoirs in boreal environments, have
 60 been shown to result in an influx of Hg to the newly
 61 created aquatic ecosystem in laboratory (Morrison and
 62 Thérien 1991), mesocosm (Hall and St. Louis 2004)
 63 and whole-lake settings (Hall et al. 2005; Bodaly et al.
 64 2007). The role of large-scale inundation events, such
 65 as those associated with impoundment for hydroelec-
 66 tric generation, are well documented sources of Hg to
 67 aquatic ecosystems (St. Louis et al. 2004), which have
 68 been shown to be present in aquatic foodwebs at
 69 multiple trophic levels (Bodaly et al. 1984; Hall et al.
 70 1998). Natural aquatic ecosystem expansion, for
 71 example due to beaver dam impoundment, has also
 72 been shown to result in enhanced mercury mobility
 73 (Roy et al. 2009). These findings have important
 74 implications for future mercury dynamics in northern
 75 landscapes, where climate change is resulting in shifts
 76 in lake hydrological regimes, including lake expansion
 77 and shoreline flooding in some instances (Carroll et al.
 78 2011; Parsekian et al. 2011).

79 A dramatic example of northern lake expansion and
 80 landscape inundation linked to recent climate change
 81 can be found on the northwest shore of Canada's Great
 82 Slave Lake, in the Great Slave Lowlands and Plains
 83 ecoregions of the Northwest Territories (Fig. 1). This
 84 landscape exhibits little relief, and has a high propor-
 85 tion of water cover, being dominated by wetlands,
 86 small ponds, and many large, shallow lake ecosystems
 87 (Ecosystem Classification Group 2009). The region
 88 also contains the Mackenzie Bison Sanctuary, estab-
 89 lished in 1963 as habitat for an ecologically important
 90 population of wood bison (*Bison bison athabascae*), a
 91 distinct sub-population, and North America's largest
 92 land mammal (Larter et al. 2000). A recent investiga-
 93 tion of changes in the Great Slave lowlands region
 94 based on Landsat satellite imagery showed that the
 95 proportion of the landscape occupied by water in a
 96 10,000 km² area (including the majority of the bison
 97 sanctuary) nearly doubled between 1986 and 2010,
 98 and that this increase was correlated with climatic
 99 variables (Korosi et al. 2017). While the whole of the

Great Slave Lowlands and Plains has become wetter 100
 recently, the response is quite heterogeneous, with 101
 some lakes exhibiting extensive (> 800%) expansion, 102
 while others showed a more muted response (Korosi 103
 et al. 2017). Lake sediment records were used to 104
 extend the record of lake area changes beyond the 105
 observational record, and showed no other periods of 106
 lake expansion as large and persistent as those 107
 occurring in the last ~ 25 years have occurred in at 108
 least two centuries (Korosi et al. 2017). 109

Landscape flooding in the Mackenzie Bison Sanc- 110
 tuary is inundating terrestrial vegetation on the 111
 margins of lakes and ponds, drowning the sedges 112
 and grasses utilized as the preferred forage by bison, 113
 and potentially driving them out of the sanctuary in 114
 search of other food sources (Korosi et al. 2017). This 115
 widespread flooding of terrestrial material may also be 116
 resulting in an influx of mercury to the lake ecosys- 117
 tems, similar to the flooding that occurs with natural 118
 impoundments, such as in beaver ponds (Roy et al. 119
 2009), and analogous to hydroelectric impoundments 120
 (Teisserenc et al. 2014). Satellite and field-based 121
 observations have shown that the recent flooding is 122
 primarily refilling old lake basins, though the lakes 123
 have not been as large as their current area in at least 124
 the last 200–300 years (Korosi et al. 2017). It is 125
 conceivable that, in this hydrologically dynamic 126
 landscape where seasonal increases in water levels 127
 may have always occurred, but the recent, persistent 128
 lake expansion is likely due to climate-related 129
 changes, flooded lakes may record a less clear 130
 response in mercury change compared to beaver 131
 impoundments and reservoirs. We explore this sce- 132
 nario using a paleolimnological approach in a strate- 133
 gically selected series of lakes in this rapidly changing 134
 region (Fig. 1). As lake sediments represent a faithful 135
 record of changes in total mercury over time (Lockhart 136
 et al. 2000), we reconstructed the recent history of 137
 mercury accumulation in order to test the hypothesis 138
 that landscape flooding in the Mackenzie Bison 139
 Sanctuary has resulted in increased mercury concen- 140
 trations in recently expanded lakes. 141

Study site 142

The study region is located within the Great Slave 143
 Plains High Boreal, and Great Slave Lowlands Mid 144
 Boreal (Level IV) ecoregions (Ecosystem Classifica- 145
 tion Group 2009). The region has limited relief, is 146

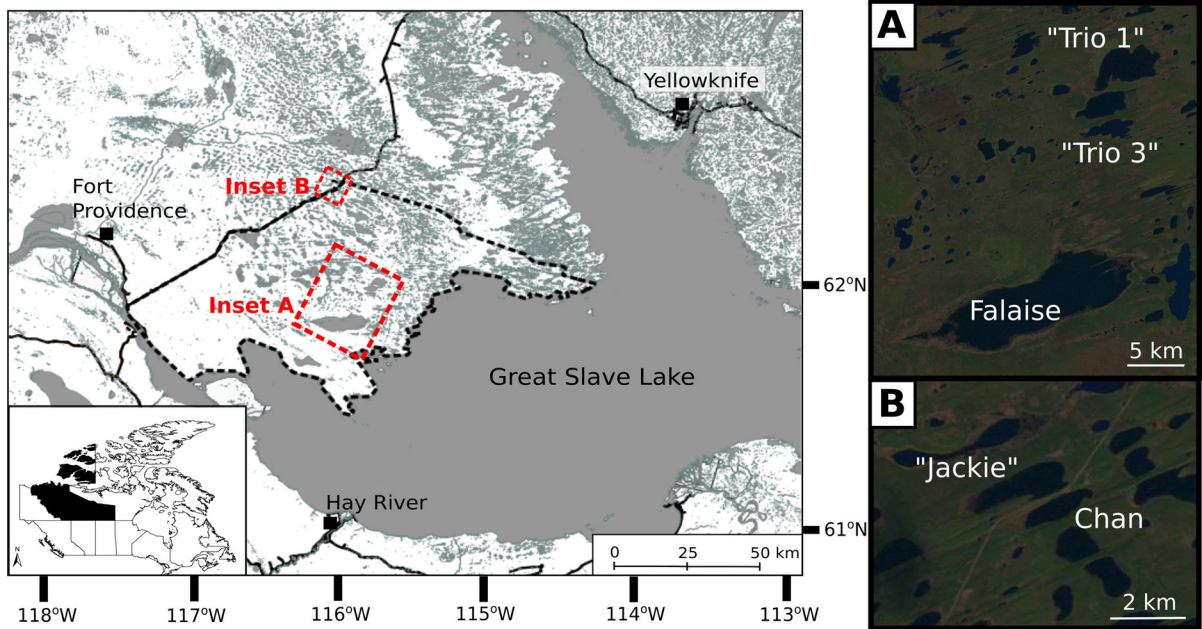


Fig. 1 (a) Map of the study area in the Great Slave Lowlands and Plains region of Canada's Northwest Territories. The area designated as the Mackenzie Bison Sanctuary is delineated with

a black dashed line. Landsat images of the study lakes from 2010 are presented in (a, b)

147 dominated by wetlands, with a multitude of small marl
 148 ponds, and a number of large, shallow lake ecosys-
 149 tems. Bedrock is primarily Cambrian to Devonian
 150 aged limestone, sandstone, and dolomite. Surficial
 151 geology is composed of a thick mantle of glacial and
 152 postglacial deposits greater than 80 m in depth (Craig
 153 1965), composed of tills and glaciolacustrine deposits
 154 from glacial Lake McConnell. Soils are variable
 155 throughout the region, including Organic soils in
 156 wetland areas, coarse-textured Brunisolic soils espe-
 157 cially near beach ridges, and Gleysols adjacent to
 158 ponds. Permafrost in the region is sporadic discontin-
 159 uous, with organic Cryosols associated with peat
 160 plateaus (Ecosystem Classification Group 2009).

161 Five study lakes were selected in the Great Slave
 162 Plains and Lowlands ecoregions for analyses (Fig. 1,
 163 Table 1). Four of the lakes are located in the core of
 164 the Mackenzie Bison Sanctuary (MBS) and to the east
 165 of NWT HWY 3, which forms the western edge of the
 166 MBS. The fifth lake, "Jackie Lake" (unofficial name)
 167 is located on the periphery of the MBS to the west of
 168 NWT HWY 3. In order to assess the potential for
 169 heterogeneous responses in mercury change associ-
 170 ated with varying lake size, the lakes range in area
 171 from the largest lake in the MBS, Falaise Lake, to

Chan Lake, a small lake located along NWT HWY 3
 (Table 1). In order to test our prediction that landscape
 flooding has resulted in increased mercury accumula-
 tion, lakes that have exhibited recent expansion are
 compared to sites that have not expanded in the recent
 past (Table 1; Korosi et al. 2017).

Materials and methods

Sediment cores were collected from the deepest
 location in each of the five lakes through the late-
 winter ice in March of 2012, using a Glew-type gravity
 corer (Glew 1989). Sediment cores were extruded into
 0.5 cm intervals using a Glew-type vertical extruder
 (Glew 1988), and kept < 10 °C during transport and
 prior to analyses. The sediment cores were the same as
 utilized for the analyses presented in Korosi et al.
 (2017), which included determination of total organic
 carbon, nitrogen content, C:N elemental ratio, and ¹³C
 and ¹⁵N stable isotope analyses. Selected intervals
 were prepared for ²¹⁰Pb and ¹³⁷Cs-based radioisotopic
 dating using gamma spectroscopy, with details pre-
 sented in Korosi et al. (2017), with the constant rate of
 supply (CRS) model used for sediment age

Table 1 Study lake locations, recent surface area and Landsat-derived area change (%) between 1986 and 2010 (from Korosi et al. 2017)

Lake	Latitude (°N)	Longitude (°W)	2010 surface area (ha)	Percent change 1986–2010 (%)
Falaise	61.47642	116.15280	5637.6	+ 824
“Trio 1”	61.64026	116.05184	1036.4	+ 462
“Trio 3”	61.59762	116.07063	306.4	+ 20
“Jackie”	61.89678	116.55987	151.0	+ 313
Chan	61.89079	116.54170	66.2	+ 51

Lake names in quotation marks are unofficial

determination (Fig. 2). Modelled errors associated with the dates corresponding to the observed recent lake expansion were low (less than 5 years for dates since 1950). Freeze-dried and homogenized samples were analyzed for total mercury by thermal decomposition with gold trap amalgamation and cold vapour atomic absorption spectrometry (CV-AAS) using a Nippon Instruments SP-3D mercury analyzer with a detection limit of 0.01 ng per sample size. Required sample masses ranged from 17 to 30 mg (dry weight). Measurement accuracy was estimated by running blanks and calibrated with MESS-3 (91 ± 9 ng g⁻¹, Natural Resources Canada) as reference material, every 10 samples. Results of mercury concentration analyses were determined per unit mass dry sediment weight, as well as per unit mass of organic carbon. As the concentration of organic carbon throughout the sediment cores did not change significantly (Korosi et al. 2017), the profiles of total mercury per unit organic carbon did not differ from per unit dry weight, and as such only the latter are discussed below. Mercury fluxes to lakes were estimated based on the modelled sedimentation rate determined via the CRS model utilized for sediment chronology development for the sediment cores for all five lakes (Fig. 2).

219 Results

220 Changes in sedimentary mercury concentrations

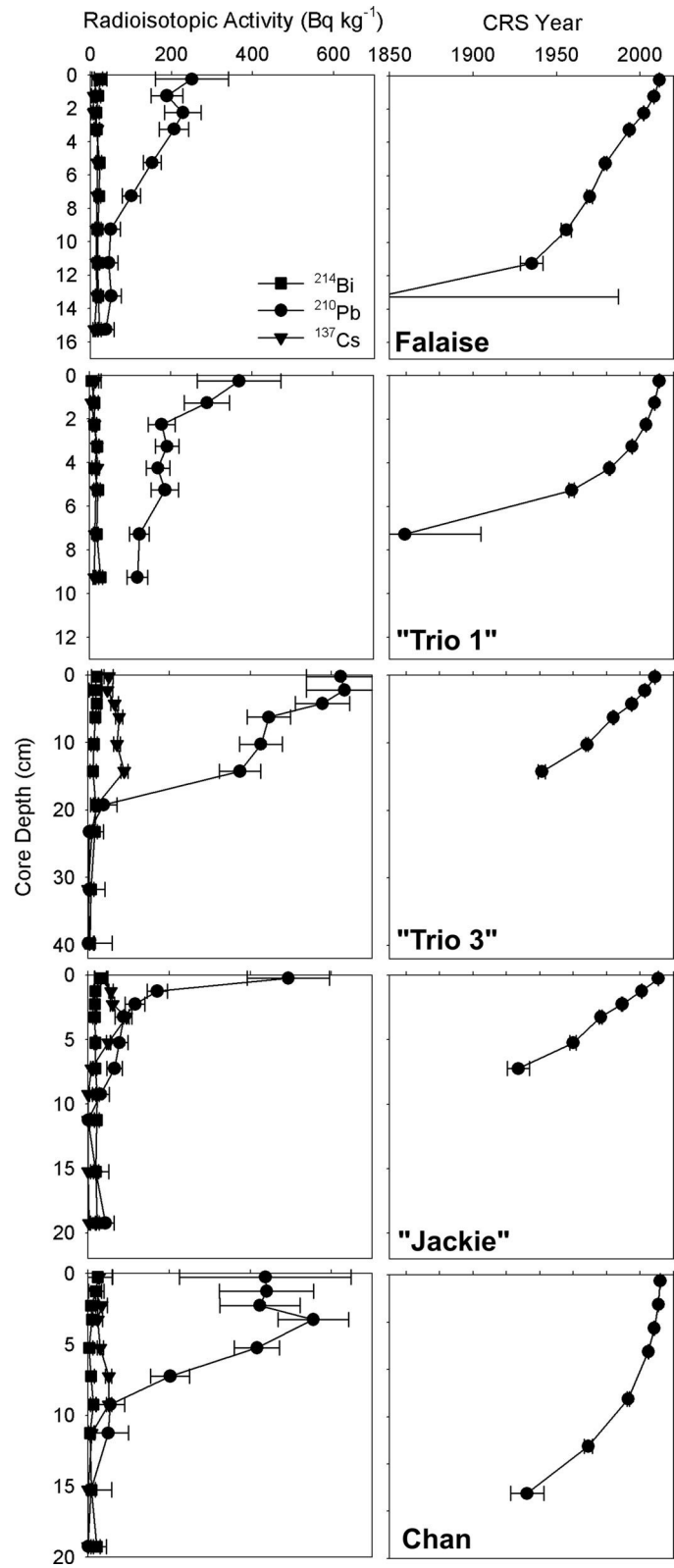
221 The concentration of total mercury (THg) in the
222 sediment core from Falaise Lake showed only small
223 variations over time (Fig. 3). Over the recent past, a
224 relatively small magnitude increase in THg was
225 observed since the mid-1980s, which was strongly

226 correlated to the marked (> 800%) increase in lake
227 area (Spearman rank correlation, $n = 7$, $r_s = 0.94$,
228 $p < 0.001$) (Fig. 3). The highest concentration
229 recorded in Falaise Lake, ~ 33 ng g⁻¹ in the surface
230 sediment interval, was well below the sediment quality
231 guidelines for the protection of aquatic life set at
232 170 ng g⁻¹ by the Canadian Council of Ministers for
233 the Environment (CCME 1999) The THg concentra-
234 tion in “Trio 1” Lake increased slightly from the
235 earliest part of the record (~ 1850) until ~ 1950 ,
236 after which it decreased, despite the increase in lake
237 surface area observed for this system (Fig. 4a). “Trio
238 3” Lake recorded the highest magnitude THg con-
239 centration in the sediment cores in this study (Fig. 4b).
240 The concentrations of THg in “Trio 3” decreased after
241 ~ 1975 , during which time the lake area exhibited
242 some fluctuation, but no directional change (Fig. 4b).
243 In “Jackie” Lake, Hg concentration exhibited a rapid
244 increase after ~ 1990 from ~ 25 to ~ 45 ng g⁻¹,
245 which tracked closely the timing and direction of lake
246 area increase inferred from Landsat imagery (Fig. 5a).
247 In Chan Lake, Hg concentration showed several small
248 magnitude changes throughout the period represented
249 by this sediment core, which were not related to lake
250 area, which increased slightly ($\sim 51\%$) over the
251 period of record (Fig. 5b). Correlations were not
252 conducted for the latter four lakes due to a low number
253 of samples that had overlapping area estimates and
254 mercury determinations. All sedimentary total mer-
255 cury values were below the CCME sediment quality
256 guidelines.

257 Changes in mercury flux

258 The total mercury flux to the sediment in Falaise Lake
259 decreased after ~ 1940 , until the onset of lake

Fig. 2 Radioisotopic activity and constant rate of supply modelled sediment age, along with associated error, for the five sediment cores from the Great Slave Lowlands and Plains region



Author Proof

Fig. 3 Sedimentary record of total mercury concentration (per gram dry sediment weight) and Landsat-derived lake surface area for Falaise Lake, Northwest Territories, Canada

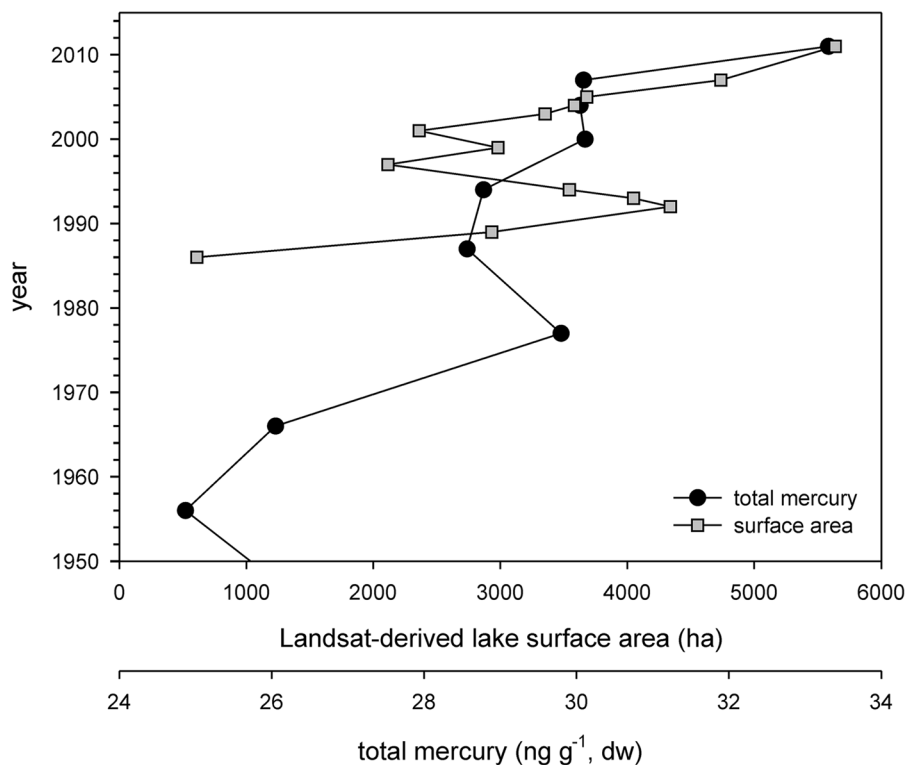
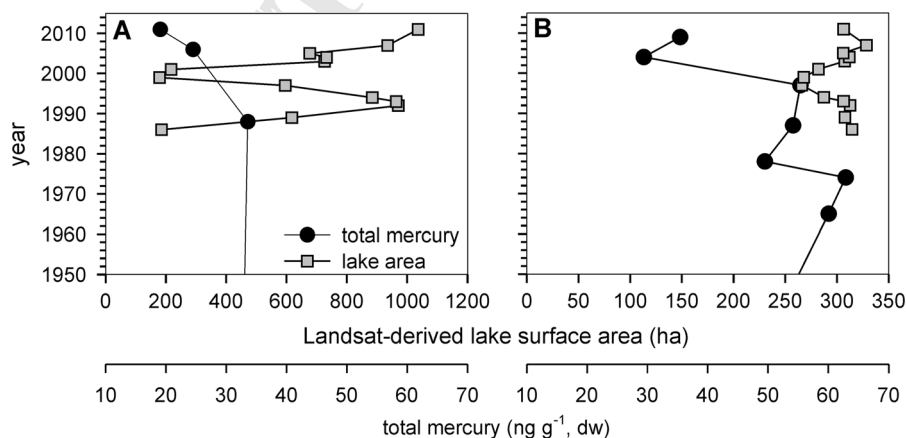


Fig. 4 Sedimentary record of total mercury concentration (per gram dry sediment weight) and Landsat-derived lake surface area for **a** Trio 1 and **b** Trio 3 lakes, Northwest Territories, Canada. The “Trio Lakes” are located close together, but have drastically different histories of recent expansion (Fig. 1, Table 1)



260 expansion in the late 1980s, after which it increased
 261 through to the top of the core (2010) (Fig. 6a). The
 262 THg flux increased from ~ 1800 ng m⁻² year⁻¹ in
 263 ~ 1980 to ~ 3000 ng m⁻² year⁻¹ in the surface
 264 sediments (Fig. 6a). Mercury flux in “Trio 1” Lake
 265 remained low throughout the record, and increased
 266 from ~ 1930 until ~ 1990, after which it was
 267 constant until the uppermost (surface) sediment inter-
 268 val, where it decreased (Fig. 6b). The sedimentation
 269 rate in Lake “Trio 3” was the lowest of the sediment

270 cores modelled in the study, and similarly THg flux
 271 was also low throughout the period for which
 272 sedimentation rate was estimated (since ~ 1960)
 273 (Fig. 6c). During this period the flux of mercury did
 274 not change markedly (Fig. 6c). Mercury flux and
 275 sedimentation rate in “Jackie” Lake increased rapidly
 276 from the bottom of the core, which continued over the
 277 recent period of rapid lake expansion (Fig. 6d). The
 278 THg flux in “Jackie” Lake was of the highest
 279 magnitude of the sediment cores from the five study

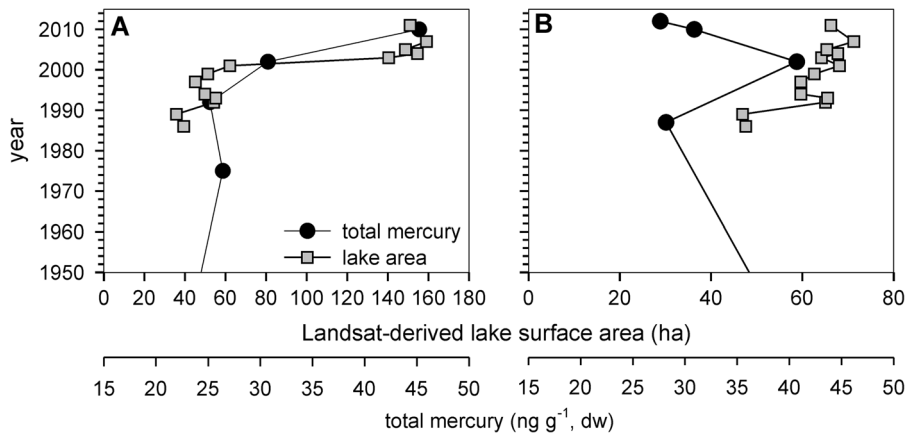


Fig. 5 Sedimentary record of total mercury concentration (per gram dry sediment weight) and Landsat-derived lake surface area for **a** “Jackie Lake” (unofficial name) and **b** Chan Lake, Northwest Territories, Canada

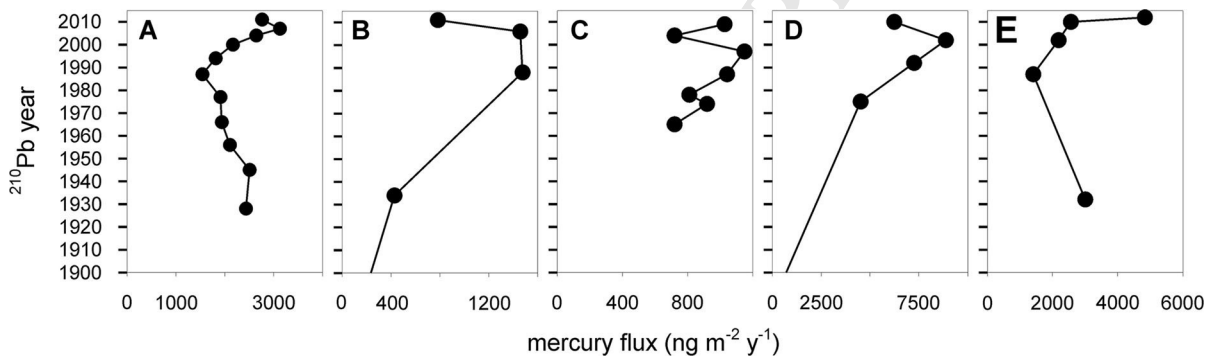


Fig. 6 Total mercury flux to the sediments (per gram dry weight, per square meter, per year) for **a** Falaise, **b** “Trio 1”, **c** “Trio 3”, **d** “Jackie”, and **e** Chan lakes, in the Mackenzie Bison Sanctuary, Northwest Territories, Canada

280 lakes, with the surface sediments recording a mercury
 281 flux of $\sim 7500 \text{ ng m}^{-2} \text{ year}^{-1}$ (Fig. 6d). In Chan
 282 Lake THg flux decreased slightly between ~ 1930
 283 and the late 1980s, after which it increased through to
 284 the surface sediment intervals, consistent with the
 285 small magnitude lake area increase (Fig. 6e). The
 286 mercury flux estimates for the smaller lakes were of a
 287 higher magnitude than the larger study lakes in the
 288 study (Fig. 6).

289 **Discussion**

290 Is landscape flooding releasing mercury to lakes?

291 The potential for mercury stored in vegetation and
 292 soils surrounding the lakes to be released following
 293 flooding has been well documented (St. Louis et al.

294 2001, 2004; Roy et al. 2009; Teisserenc et al. 2014). 294
 295 However, most studies on mercury release in larger- 295
 296 scale lake ecosystems have been due to anthropogenic 296
 297 impoundment for hydroelectric generation, and not 297
 298 due to natural landscape flooding. Natural impound- 298
 299 ment, such as through the bio-engineering activities of 299
 300 beavers, are also a source of mercury to aquatic 300
 301 ecosystems (Roy et al. 2009), though the magnitude of 301
 302 impoundment from these actions is smaller. The recent 302
 303 lake area changes of Falaise Lake are correlated with 303
 304 increased total mercury, though the magnitude of the 304
 305 increase is small (Fig. 3). Similarly, total mercury in 305
 306 “Jackie” Lake, a much smaller ecosystem, has 306
 307 increased coincident with recent lake expansion 307
 308 (Fig. 5a). In “Trio 1” Lake, which has undergone 308
 309 the second greatest increase in surface area of the 309
 310 study lakes, total mercury has decreased over the 310
 311 recent past (Fig. 4a). The highest overall 311

concentration of total mercury in the sediments of the five study lakes was recorded in Lake “Trio 3” which has not expanded recently (Fig. 4b). In Chan Lake, which has increased in surface area moderately compared to the other lakes (51%), there was no concurrent, directional increase in total mercury concentration in the sediments (Fig. 5b). Sedimentary mercury concentrations found throughout the profiles of all five sediment cores analyzed in this study are well below the sediment quality guidelines for the protection of aquatic life set at 170 ng g^{-1} by the Canadian Council of Ministers for the Environment (CCME 1999), and the range of values are similar to those found in other studies of subarctic lakes in Canada (Lockhart et al. 1995, 1998; Muir et al. 2009; Brazeau et al. 2013).

Increases in fluxes of mercury to the sediment correspond relatively closely to the timing of the recent lake expansion in Falaise and “Jackie” lakes, as well as in Chan Lake, which increased marginally in area. Despite these increasing mercury flux trends, the absolute values recorded are low, at least an order of magnitude less than lakes in northern Scotland, for example (Yang and Rose 2003). The highest flux values (Chan and “Jackie” lakes) were similar to the lowest values documented for undisturbed sites from northern Canada (Lockhart et al. 1995), western Greenland (Bindler et al. 2001) and Scandinavia (Renberg 1986; Verta et al. 1989). Both the mercury concentration and flux data suggest that landscape flooding in the Great Slave Lowlands region has not resulted in the release of large amounts of total mercury to the sediments of impacted lake ecosystems.

The variable trend of mercury change, as well as the small magnitude of increase in both concentration and flux in those sites that are recording increases, contrasts with the well-documented enrichment of mercury in lake water and sediments following reservoir impoundment (Teisserenc et al. 2014). There are several reasons why this may be the case. Though the persistence of recent expansion is unique over the last several 100 years, these shallow basins are undoubtedly wetted periodically during snowmelt or periods of heavy precipitation. This would result in relatively consistent and regular leaching of mercury into the lake, as opposed to soils in reservoir impoundments that have been building up inorganic mercury over long timescales, which are then

incorporated into lakewater and sediment rapidly, and in bulk. In addition, the composition of the soils could result in conversion to gaseous Hg^0 , which would be lost to the atmosphere, and not transported to the lake sediments (Poulin et al. 2016), and thus detailed sampling of local soil properties would be helpful for understanding mobilization and transformation properties. A second confounding factor may be the regular occurrence of fire in this region, which burns soil organic matter reserves and depletes soil mercury concentrations, also through release of gaseous mercury (Mailman and Bodaly 2005; Friedli et al. 2003). The burning of land prior to impoundment has been suggested as a mechanism for decreasing methylmercury production, through the loss of inorganic mercury (Mailman et al. 2006). Fire regime is a critically important control on vegetation structure within subarctic boreal regions (Johnson 1979), and its influence may extend to landscape mercury cycling as well. Fire return intervals are projected to decrease in the Great Slave region as a result of climate warming (Boulanger et al. 2014), and historical fire frequency may be higher in the Northwest Territories region than in eastern Canada where most of the studies of impoundment-driven mercury enrichment have occurred (Bergeron et al. 2004).

Total organic content (TOC) in the sediments of all of the lakes is relatively low $\sim 20\text{--}30\%$, and did not change as a result of recent flooding (Korosi et al. 2017). Non-peat soils in the Great Slave Lowlands region, derived from Glacial Lake McConnell sediments (Smith 1994) tend to have a high proportion of clay and are relatively low in organic carbon (Ecosystem Classification Group 2009), in comparison to nearby areas such as the Tathlina and Kakisa lake watersheds to the west and south (TOC averages $40\text{--}50\%$), where strong increases in mercury have been reported (Korosi et al. 2015). The concentrations of mercury in uncontaminated soil samples were measured at 55 ng g^{-1} in both the B and C horizons of brunisolic samples from near Fort Providence, less than 70 km from Falaise Lake (McKeague and Kloosterman 1974). This suggests Great Slave Plains and Lowlands soils derived from Glacial Lake McConnell sediments are likely naturally low in mercury. Importantly, the local distribution of sporadic permafrost in the region is not well documented. Frozen ground may play a role in influencing the movement of mercury in soils, by limiting mobility,

410 and future thaw could result in altered mercury
411 availability, especially during fire events (Turetsky
412 et al. 2006).

413 It is important to note that throughout this study we
414 have focused on reconstructing the history of total
415 mercury changes in these lake ecosystems through
416 sediment records. We have not reconstructed changes
417 in methylmercury, the most toxic and bioaccumulative
418 form of mercury, as historical trends cannot be reliably
419 reconstructed from lake sediment records. There exists
420 the potential for landscape flooding in the region to
421 alter rates of methylation and demethylation, partic-
422 ularly if peatlands are inundated (Heyes et al. 2000),
423 and increased deposition of mercury is not necessarily
424 required for methylmercury to enter and accumulate in
425 foodwebs (Bodaly and Fudge 1999). As changing
426 methylation activity would not be obvious from
427 sedimentary profiles of total mercury, further study
428 is needed before we can definitively say that expansion
429 has not altered the mercury cycling and availability in
430 these lakes.

431 Conclusions

432 The recent, extensive landscape flooding that has
433 occurred in the Great Slave Lowlands, including the
434 Mackenzie Bison Sanctuary, resulted in increases in
435 total mercury and mercury flux in the sediments from
436 two recently flood lakes, though the magnitude of
437 these increases was small ($\sim 25\text{--}33 \text{ ng g}^{-1}$ in Falaise
438 Lake, and $\sim 25\text{--}45 \text{ ng g}^{-1}$ in “Jackie” Lake). In
439 another flooded site, no increase in total mercury was
440 observed. Total mercury concentrations in sediments
441 in all of the lakes studied in the region are well below
442 Canada’s sediment quality guidelines, set at
443 170 ng g^{-1} . The potential for landscape flooding to
444 have resulted in changes in mercury cycling, including
445 altered rates of methylation, remains an important
446 knowledge gap that requires future study.

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