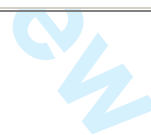




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Insights into convective momentum transport and its parametrisation from idealised simulations of organised convection

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Deep convection is a multiscale process that significantly influences the budgets of heat, moisture and momentum. In global climate models the thermodynamic effects of convection are normally treated by parametrisation schemes, with a separate formulation for convective momentum transport (CMT). The transport modules for current thermodynamic and momentum parametrisations are upright entraining plume models that do not account for vertically tilted mesoscale circulations that characterise organised convection in sheared environments. The associated countergradient vertical transport of horizontal momentum fundamentally affects dynamical interactions between the convection and the mean flow. This study examines the CMT properties of simulated idealised mesoscale convective systems, including their sensitivity to horizontal resolution, domain size, and lateral boundary conditions. It is found that even for large domains, the horizontal gradient terms are important, especially the mesoscale pressure gradients that are neglected in CMT parametrisations. A nonlinear analytic model provides a dynamical foundation for the effects of convective organisation, including the role of the horizontal pressure gradient. It is found that a small computational domain adversely affects the convective organisation by generating artificially large compensating subsidence and an unrealistic evolution of the CMT. Finally, analyses of the cross-updraft/downdraft pressure gradients expose significant uncertainties in their representation in contemporary CMT parametrisation schemes.

Key Words: convective momentum transport; mesoscale convective systems; parametrisation; Reynolds stress

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1. Introduction

Deep convection is a globally important process that redistributes heat, moisture and momentum vertically throughout the troposphere. Convective momentum transport (CMT) is known to have a considerable effect on the both the zonal and meridional mean flow (Zhang and McFarlane (1995)). Earlier studies have shown it contributes to the Hadley circulation (Schneider (1975), Helfand (1979)), with Houze (1973) noting that CMT can have a similar magnitude to the rest of the terms comprising the atmospheric angular momentum budget. CMT is normally parametrised in global models, but uncertainties remain about the dynamics of CMT, its effects, and its representation in current schemes. This study revisits the problem of CMT using idealised simulations to assess some of the assumptions underlying its parametrisation.

CMT occurs on multiple scales, both temporally and spatially. Convective-scale transports are caused by unorganised convection, or the convective updrafts embedded within organised systems. The archetypal model described by Moncrieff (1992)

demonstrates how the sign of mesoscale momentum transports by organised convection are inherently linked to the tilt of the mesoscale convective system (MCS). His Fig. 10 shows that a downshear-tilted circulation produces correlated u and w terms resulting in positive CMT, and an upshear-tilted circulation produces anti-correlated terms, resulting in a negative CMT. The latter is salient herein. These countergradient mesoscale fluxes strongly affect the mean flow, feedback to the convection itself (i.e. Moncrieff (1981), Flatau and Stevens (1987), Mechem et al. (2006), Mahoney et al. (2009)), and contribute to a variety of convective regimes (Lane and Moncrieff (2010)). LeMone (1983) analysed the effect of momentum transports across and along observed convective lines and found that organised convection indeed transports zonal momentum either upgradient or downgradient, thereby enhancing or reducing the vertical wind shear accordingly. In other words, organised mesoscale CMT can defy mixing-length theory (i.e., the representation of CMT as a diffusion process) in situations where countergradient momentum

transport occurs. Depending on its characteristics, CMT can either strengthen a convective system and increase its longevity, or diminish it. Rotunno et al. (1988) and Weisman and Rotunno (2004) explain why a system changes from downshear tilted to upshear tilted as it transitions from its initial stages to a mature system. It follows that the sign of the momentum transport can change from positive to negative throughout the MCS lifecycle.

Convective-scale momentum transports have been parametrised based on entraining plume models (e.g. Kershaw and Gregory (1997)) using the convective mass flux from the convective parametrisation scheme and empirical relationships that relate the cross-updraft pressure gradient to the large-scale vertical wind shear. Application of CMT parametrisations to global models have shown demonstrable improvement in the large-scale circulations including surface winds and precipitation (Richter and Rasch (2008)). Tropical circulations such as the Hadley cell, ITCZ, MJO and ENSO (Zhang and McFarlane (1995), Mapes and Wu (2001), Wu et al. (2003), Miyakawa et al. (2012)) are also better represented when CMT is accounted for. For example, Han and Pan (2006) used the Wu and Yanai (1994) parametrisation scheme, which reduced the forecast track error of hurricane prediction and improved the intensity. Parametrisation of CMT within regional-scale models has also been shown to affect the motion of MCSs (Mahoney et al. (2009)). Further highlighting the importance of CMT, Moncrieff and Liu (2006) demonstrated that a regional numerical weather prediction (NWP) model without parametrised CMT formed mesoscale convective systems with erroneous tilts and resolved CMT of the wrong sign (when compared to a higher resolution convection-permitting model).

A few different approaches have been used to parametrise CMT. The scheme developed by Schneider and Lindzen (1976) was based upon the work by Ooyama (1971) using a mass-flux approach to cumulus parametrisation and the assumption that in-cloud horizontal momentum is conserved. However, this assumption is only valid if a negligible horizontal across-cloud pressure gradient is present. LeMone (1983) determined that across-cloud pressure gradients have an important role in influencing in-cloud wind speeds which, along with observational work by Shapiro and Stevens (1980) led to the inclusion of the effects of pressure gradients in the Flatau and Stevens (1987) parametrisation scheme. Zhang and Cho (1991) also included pressure gradients by employing a simple model of flow around a cloud. Gregory et al. (1997) include parametrised across-updraft and downdraft pressure gradients to determine in-cloud horizontal velocities; this parametrisation scheme defines the pressure gradient across the updrafts to be proportional to the mass flux and the vertical wind shear. This relationship introduces a ‘tunable’ parameter to the Gregory et al. (hereafter referred to as GKI) scheme, which has been examined by Grubišić and Moncrieff (2000) and will be considered later in this study. It is apparent that the representation of across-updraft and downdraft horizontal pressure gradients have emerged as one of the key uncertainties in the parametrisation of CMT. However, recent modelling work by Romps (2012) has questioned its importance for unorganised convection. This issue is revisited in section 3.4. Nevertheless, it is critical to note that all of the above parametrisation schemes use simple plume models to represent the convection and therefore are unable to fully represent the transports associated with organised circulations, viz. mesoscale momentum transports.

Recent work by Moncrieff et al. (2017) uses a multiscale coherent structure parameterisation (MCSP) which incorporates the effects of organised convection that are currently missing from current parameterisations. Other studies (e.g. ?) use the classification of convection by Johnson et al. (1999) in a multicloud model (MCM). This research reinforces previous work

that the role of vertical shear is important in producing the appropriate rearward slant of organised systems (e.g. Wu and Yanai (1994), Richter and Rasch (2008), although as Moncrieff et al. (2017) notes, it plays a passive role). The use of MCSP as a parametrisation for organised convection essentially accounts for the heating and momentum transport by organised convection, in cumulus parametrisation. This work supports the findings from this study.

A number of authors have used explicit models of moist convection to study CMT (e.g., Grubišić and Moncrieff (2000), Gao et al. (1990), Gallus and Johnson (1992) and Yang and Houze (1996)) and the underlying dynamics. In our study we build on this previous work by using a variety of idealised simulations to test some of the assumptions inherent in the CMT parametrisation schemes. Specifically, a range of model domain sizes are used to assess the contributions of terms involving horizontal gradients, which are normally neglected in parametrisations. Their sensitivity to model resolution is also considered. The across-updraft and downdraft pressure gradients are examined in the context of the relationships used in the GKI parametrisation and to assess the relative contributions of the convective-scale transports to the domain-mean tendency.

The remainder of the paper is ordered as follows: the numerical model and its configuration is described in section 2, along with the decomposition of the momentum budget. The results of the idealised simulations are presented and analysed in section 3, along with an evaluation of Moncrieff’s (1992) analytic model. The paper is summarised in section 4.

2. Idealised model simulations

2.1. Model description and configuration

All simulations use the Weather Research and Forecasting Model (WRF) - Advanced Research core (ARW) version 3.3 (Skamarock et al. (2008)). WRF-ARW is a mesoscale modelling system, developed principally by the National Center for Atmospheric Research and is used for real-time and operational forecasting as well as research applications. The model uses the fully compressible Eulerian nonhydrostatic equations, which are solved on a finite-difference mesh with Arakawa C-grid staggering and formulated using a mass-based vertical coordinate (Laprise (1992)). In this study moist processes are treated explicitly (no cumulus parametrisation) and represented by the WRF single-moment 6-class microphysics (WSM6) scheme (Hong and Lim (2006)). The effects of subgrid turbulence are parametrised using a predictive 1.5-order turbulence kinetic energy closure. For this study Coriolis effects and other physical processes (viz. radiation, surface friction, surface fluxes) are neglected.

The height of the computational domain is 35 km and a 10 km Rayleigh damping layer is imposed to prevent gravity wave reflection from the upper boundary. Three computational domains are used: 100 km x 50 km (denoted small, Sm), 400 km x 200 km (denoted medium, Med) and 800 km x 400 km (denoted large, Lg). Another domain - 50 km x 50 km (denoted square, Sq) was also run with 1 km horizontal grid spacing (Sq_1km) only. For each of the main three domain sizes, three different horizontal grid spacings were used to explore resolution sensitivities (500 m, 1 km and 3 km). There are 71 vertical levels with an approximate vertical grid spacing of 500 m for all domains. The timesteps are 3 s, 6 s and 12 s for 500 m, 1 km and 3 km gridlengths, respectively. All model simulations are run for a duration of 6 hours - the domain configurations are summarised in Table 1. All domain configurations are run for both open and cyclic boundary conditions, except the Sq domain which uses cyclic boundary conditions only.

Table 1. Simulation domain configurations.

Model run	Domain size	$\Delta x, \Delta y$
Sq_1km	50 km x 50 km	1 km
Sm_500m	100 km x 50 km	500 m
Sm_1km	100 km x 50 km	1 km
Sm_3km	100 km x 50 km	3 km
Med_500m	400 km x 200 km	500 m
Med_1km	400 km x 200 km	1 km
Med_3km	400 km x 200 km	3 km
Lg_500m	800 km x 400 km	500 m
Lg_1km	800 km x 400 km	1 km
Lg_3km	800 km x 400 km	3 km

The initial model thermodynamic environment for all simulations is defined using the Weisman and Klemp (1982) analytic sounding (Fig 1(b)-(c)). The initial wind profile is shown in Fig. 1(a) and contains constant vertical shear near the surface such that

$$U_0(z) = \begin{cases} 0, & z > h; \\ U_0 - \frac{U_0 z}{h}, & z \leq h. \end{cases} \quad (1)$$

where $U_0 = -18.0 \text{ m s}^{-1}$ and $h = 3.25 \text{ km}$. This wind profile is used because previous studies have shown that this low-level linear shear produces a strong, long-lived MCS (Thorpe et al. (1982), Weisman and Rotunno (2004)). The meridional wind component v is initially zero. Convection is initiated by a temperature perturbation ('warm bubble') of 3 K in the horizontal centre of the domain with a radius of 10 km and a height of 1.5 km. The mean wind profile is free to evolve with the simulation; an example evolution of the mean zonal wind u is shown in Fig. 1(a) (from Med_1km, See Table 1, with cyclic boundary conditions). This evolution represents the effect of the mean momentum tendency, which leads to an increase in surface u , a decrease in upper-level u and an eventual reduction in the low-level shear.

2.2. Convective system

The result of all simulations is a three-dimensional mesoscale convective system of the leading convective-line trailing-stratiform type (e.g. Houze et al. (1989), Parker and Johnson (2000)), though there are structural differences in spatial extent and strength between the convection in the various size and resolution domains. As an example, the evolution and mature structure of a simulated storm (Med_1km, see Table 1) is shown in Figs. 2 & 3. The system covers about 200 km and 100 km in the x- and y-directions respectively (Fig. 3). The vertical cross-section of the zonal wind in Fig. 2(c) implies the triple-branch model from Moncrieff (1992) (his Fig. 1(b)) with front-to-rear ascending flow, an overturning updraft branch, and a descending mesoscale downdraft at the rear of the system. A stagnation point is present at the surface, near $x=170 \text{ km}$, identifying the leading edge of the cold pool that helps maintain the system by triggering new convection. The contours in Fig. 2(d) are perturbation pressure and the red contours represent the area of greatest negative perturbation pressure, which indicates the position of the mesolow, located behind the updraft. This is reinforced by the presence of a region of higher pressure at the front of the system.

The mean wind profile has positive low-level wind shear (Fig. 1) and Fig. 2 demonstrates that the main convective region of the mature system is upshear-tilted. The simulated systems evolve in a similar way to that depicted by Weisman and Rotunno (2004), with the convective region being downshear-tilted in its earliest stages and becoming upright as the cold pool develops (Fig. 2(b)). This evolution will be considered later.

2.3. The momentum budget

The form of the x-component of the inviscid horizontal momentum equations in height coordinates are:

$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} + w \frac{\partial u}{\partial z} = -\frac{1}{\rho} \frac{\partial p}{\partial x}, \quad (2)$$

where, u , v and w are the three velocity components, p is pressure, ρ is density, and t is time. As WRF is a fully compressible numerical model, the mass continuity equation is:

$$\frac{\partial \rho}{\partial t} + \frac{\partial \rho u}{\partial x} + \frac{\partial \rho v}{\partial y} + \frac{\partial \rho w}{\partial z} = 0 \quad (3)$$

Each of the variables are separated into the temporally varying horizontal domain-mean wind and corresponding perturbation. The variables are now decomposed as $u = \bar{u}(z, t) + u'$, $v = \bar{v}(z, t) + v'$, and $w = \bar{w}(z, t) + w'$, where the overbar represents the horizontal domain average at each time and the primes denote perturbations. For the zonal wind component this becomes:

$$\begin{aligned} \frac{\partial \bar{u}}{\partial t} + \frac{\partial u'}{\partial t} + \frac{\bar{u}}{\rho} \frac{\partial \rho}{\partial t} + \frac{u'}{\rho} \frac{\partial \rho}{\partial t} + \frac{1}{\rho} \frac{\partial \rho \bar{u}}{\partial x} + \frac{2}{\rho} \frac{\partial \rho \bar{u} u'}{\partial x} + \frac{1}{\rho} \frac{\partial \rho u' u'}{\partial x} + \\ \frac{1}{\rho} \frac{\partial \rho \bar{u} v}{\partial y} + \frac{1}{\rho} \frac{\partial \rho \bar{u} v'}{\partial y} + \frac{1}{\rho} \frac{\partial \rho u' v}{\partial y} + \frac{1}{\rho} \frac{\partial \rho u' v'}{\partial y} + \frac{1}{\rho} \frac{\partial \rho \bar{u} w}{\partial z} + \\ \frac{1}{\rho} \frac{\partial \rho \bar{u} w'}{\partial z} + \frac{1}{\rho} \frac{\partial \rho u' w}{\partial z} + \frac{1}{\rho} \frac{\partial \rho u' w'}{\partial z} = -\frac{1}{\rho} \frac{\partial p}{\partial x} \end{aligned} \quad (4)$$

Each term in the above equation is then averaged horizontally across the domain (i.e., Reynolds averaging). After rearranging, removing terms that are identically zero, assuming $\bar{v} = 0$, and neglecting a few small terms this becomes:

$$\begin{aligned} \frac{\partial \bar{u}}{\partial t} = -\frac{\bar{u}}{\rho} \frac{\partial \rho u'}{\partial x} - \bar{w} \frac{\partial \bar{u}}{\partial z} - \frac{1}{\rho} \frac{\partial \rho u'^2}{\partial x} \\ - \frac{1}{\rho} \frac{\partial \rho u' v'}{\partial y} - \frac{1}{\rho} \frac{\partial \rho u' w'}{\partial z} - \frac{1}{\rho} \frac{\partial p}{\partial x} \end{aligned} \quad (5)$$

where each term now represents a horizontal mean, and variables with overbars ($\bar{\quad}$) represent the domain-mean value of the variable at each height level. Note here that the full pressure is retained for simplicity, however as the pressure term involves only a horizontal derivative it is exactly equivalent to using a pressure perturbation from the horizontal mean. Note that for this paper, the equation for v (the meridional wind component) is not considered as there is negligible mean tendency on v because the mean wind and shear are aligned with the x-direction.

For periodic lateral boundary conditions, the mean vertical velocity (\bar{w}) and all of the terms involving horizontal derivatives are exactly zero. Thus, this simplification of Eqn. 5 leads to the well-known result:

$$\frac{\partial \bar{u}}{\partial t} = -\frac{1}{\rho} \frac{\partial \rho u' w'}{\partial z} \quad (6)$$

This infers that, for periodic conditions, the total tendency imposed on the mean flow can be wholly represented by the momentum flux divergence term $\frac{1}{\rho} \frac{\partial \rho u' w'}{\partial z}$. This result forms the basis of all parametrisation schemes of CMT.

3. Results

To calculate the momentum budgets and tendencies, the model output was first interpolated from the WRF native mass-based

coordinate system to height coordinates. This interpolation was necessary to close the momentum budget as written in Eq. 5. The sum of all the terms comprising the momentum budget (Eq. 5) along with the mean tendency of \bar{u} , were calculated for all simulations with the difference between the sum of the terms and the mean tendency defining a residual. Budgets were initially compared for calculations using model output every 3, 15, 30 and 60 minutes for the Med_1km_open simulation (not shown). As expected, the budgets with the 3-minute output showed the smallest residual compared to the other time resolutions; 15-minute data provided similar results, with a notable increase in the residual for 30- and 60-minute data. The small amount of added accuracy for the 3-minute output was not worth the considerably larger data storage, and 15-minute model output is used for all budgets found herein.

3.1. Analysis of momentum budgets and tendencies

In order to compare the domain-mean vertical flux of horizontal momentum with the other terms comprising the momentum budget (Eqn. 5), Fig. 4(a) shows all terms that contribute to the mean tendency for the open boundary conditions, medium domain. It is obvious that although the momentum flux ($\rho u' w'$) term dominates the other terms and accounts for the overall sign of the tendency across the domain in the upper-levels, there are other terms that significantly contribute to the tendency on the mean flow. The terms that are most important within the momentum budget are $\frac{1}{\rho} \frac{\partial \rho u' w'}{\partial z}$, $\frac{1}{\rho} \frac{\partial \rho u' u'}{\partial x}$ and $\frac{1}{\rho} \frac{\partial p}{\partial x}$. As stated, the sign of the total tendency (indicated by the solid red line) at upper levels is consistent with the divergence of the vertical flux of horizontal momentum but this is offset by $\frac{1}{\rho} \frac{\partial p}{\partial x}$ and $\frac{1}{\rho} \frac{\partial \rho u' u'}{\partial x}$, which comprise the static and dynamic Bernoulli pressure terms, respectively.

The sum of all the budget terms (Fig. 4(b) - red line) is very similar to the actual tendency, i.e., there is a very small residual. The largest residual is found near the tropopause at approximately 10.5 km altitude. This larger residual at these heights is related to transient behaviour, presumably linked to overshooting convective updrafts, because calculations with 3-minute data showed a smaller residual at those heights.

In the first two hours of the simulation the mean momentum flux and tendency is much smaller than at other times (and over the whole 6-hour simulation) (Fig. 4(c)). The tendency has opposite sign at early times near the tropopause compared to later times, but is very small (Fig. 4(d)). Fig. 2(b) shows a slight downshear tilt suggesting that any transports due to mesoscale organisation, albeit weak, are potentially offset by those from convective-scale transports as the domain-mean momentum flux is near zero.

The three largest terms - $\frac{1}{\rho} \frac{\partial \rho u' w'}{\partial z}$, $\frac{1}{\rho} \frac{\partial \rho u' u'}{\partial x}$ and $\frac{1}{\rho} \frac{\partial p}{\partial x}$ are also averaged across the convective updrafts and downdrafts to determine the convective-scale behaviour and its contributions to the domain-mean momentum budgets. Here updrafts are defined as regions within cloud that have vertical velocity greater than 1 m s^{-1} , and downdrafts are regions within cloud with vertical velocity less than -1 m s^{-1} . Figure 5 shows how these small-scale contributions to the tendency change as the convective system evolves. During the initial stages (i.e. the first two hours), the pressure gradient across updrafts is positive, changing to negative at low levels during the later stages as the system becomes organised. This suggests the pressure gradient acts to accelerate the low level flow in the downshear direction in the early stages and in the upshear direction in the mature stages. At low-levels the magnitude of this term increases as the system develops. The pressure gradients across the downdrafts generally show a weaker tendency of the opposite sign to the updraft tendencies.

The total cross updraft/downdraft momentum flux $\rho u' w'$ divergence (Fig. 5(c)) increases in magnitude as the system matures. The tendency across the updrafts acts to increase the zonal flow below about 4 km and decrease the flow above 4 km, consistent with net negative momentum flux and downgradient transport. The momentum flux tendency associated with the updrafts is larger than that from the downdrafts, which is mostly confined below 4 km. This result is consistent with intense convective cores and evaporatively driven downdrafts, which act to partially offset the tendency from updrafts.

Weisman and Rotunno (2004) explain how mesoscale convective systems evolve from being downshear tilted to upshear tilted as the baroclinically generated vorticity from the cold pool begins to dominate the environmental shear at low-levels. This explanation is consistent with the evolution of the convective system herein, with (Fig. 2 (a)-(c)) showing the transition from downshear to upright to upshear tilted over about 3 hours. However, the tendency from the updraft/downdraft momentum flux divergence (Fig. 5(c)) does not undergo this transition and maintains its vertical structure throughout the evolution. The only significant change from the early stages to the mature stages is the change in the cross-updraft pressure gradient. Later in the storm evolution the upshear-tilted mesoscale circulations work in concert with the convective-scale transports.

An interesting point to note is that the $\frac{1}{\rho} \frac{\partial \rho u'^2}{\partial x}$ term (Fig. 5(b)) across up/downdrafts, reinforces the momentum flux term (i.e. strengthens the tendency) at low altitudes. In the domain mean, however, the low altitude tendency from this term is small, which suggests the domain-mean mesoscale contribution balances the convective-scale contribution. Moreover, the total up/downdraft contributions from the three largest terms (Fig. 5(d)) are consistent in sign with the domain-mean contributions at low altitudes ($< 4 \text{ km}$). However above this level, the sign is opposite to the domain-mean terms in Figure 4(a). This difference in sign above 4 km suggests that the mesoscale tendencies oppose and dominate the convective-scale tendencies above the background shear layer.

To demonstrate the convective and mesoscale contributions more clearly, in Fig. 6 the contributions from the updrafts and downdrafts to $\frac{1}{\rho} \frac{\partial \rho u' w'}{\partial z}$, $\frac{1}{\rho} \frac{\partial \rho u' u'}{\partial x}$ and $\frac{1}{\rho} \frac{\partial p}{\partial x}$ are averaged across the entire domain (instead of only the updraft and downdraft areas, as in Fig. 5). The difference between the net tendency (shown in red in Fig. 6) and the convective (up/downdraft) contributions from each of these terms represents the mesoscale tendencies (here denoted 'other'). It is clear from Fig. 6 that the mesoscale tendencies play an important role in offsetting the convective tendencies, especially the $\frac{1}{\rho} \frac{\partial \rho u'^2}{\partial x}$ term (Fig. 6b) and the sum of these three largest terms (Fig. 6d).

3.2. Effect of domain size on momentum budget

Figure 7 compares the momentum budgets for the medium and large domains. For the larger domain the mean tendencies are smaller, due to the larger averaging area. As expected, the horizontal derivatives becoming less important the farther the domain boundary is from the convective area and tending to zero with increasing domain size. For the large domain $\frac{1}{\rho} \frac{\partial \rho u'^2}{\partial x}$ is negligible. However, an important feature of Fig. 7(c) is that even in the large domain, the domain-mean pressure gradient is still an important component of the momentum budget. The pressure gradient generally reduces the magnitude of the tendency caused by the momentum flux, and the upper parts of this profile are consistent with the mesoscale pressure gradient identified in Fig 6(a). Thus, the mesoscale pressure gradient is an influential component of the momentum budget, even on scales much larger than the convective system, and acts to maintain the organised mesoscale circulation. It is important to note that this

pressure gradient term is fundamentally different to the large-scale pressure gradient that controls the evolution of the large-scale flow. This mesoscale pressure gradient (and its convective-scale counterparts) are produced entirely by the convective system, and appear to be an important part of the convection upscale growth process. The relevance of this result is that current CMT parametrisation schemes neglect this mesoscale pressure gradient term and only attempt to represent the impact of the convective-scale pressure gradient term on the in-cloud velocities (an example is the GKI scheme).

The tendencies for the cyclic simulations, however, are comprised solely of the momentum flux divergence ($\frac{1}{\rho} \frac{\partial \rho u' w'}{\partial z}$), as is assumed for all parametrisation schemes. This is because the horizontal derivatives - $\frac{1}{\rho} \frac{\partial p}{\partial x}$ and $\frac{1}{\rho} \frac{\partial \rho u'^2}{\partial x}$ - are identically zero. Of interest is that for each domain size, the \bar{u} tendency for the domains with open boundary conditions (Fig. 7 (left column)) are similar to those for cyclic boundary conditions (right column). This is perhaps not surprising for the large domain as the horizontal gradient terms are small (see discussion above). However, this result is not necessarily expected for the medium domain. In particular, for the open boundary conditions the tendency from the momentum flux is large in magnitude and offset by the horizontal gradient terms. For the cyclic boundary conditions the momentum flux tendency (which is equal to the total tendency) is much smaller, allowing the \bar{u} tendency for each simulation to be very similar. It might be as the net effect of the cyclic boundary conditions (and stronger descent, see discussion later) is to reduce the strength of the convection and hence the momentum flux, with the agreement in \bar{u} tendency being fortuitous.

3.3. Resolution sensitivity

The momentum budgets for the different horizontal resolutions are compared (Fig. 8) and it is evident that the resolution does not greatly affect the relative contributions from the terms that comprise the momentum budgets. Most of the terms from Eqn. 5 contribute to the total tendency for the medium domain. The strength of the domain mean momentum flux and pressure gradient tendencies in the upper troposphere is largest for the 3km model (Fig. 8(c)) compared to the finer grid spacing simulations. At lower altitudes the differences are smaller. This upper-level sensitivity is likely because there is less entrainment and mixing in the 3 km simulation and therefore there is greater vertical transport of horizontal momentum beyond the middle troposphere.

Direct comparison of the domain mean tendencies (Fig. 9(a)) show the largest differences above the low-level shear layer in regions where convective updrafts and mesoscale circulations are strong. The 3 km grid spacing simulation shows a region of positive tendency at about 12 km, which is likely related to enhanced overshooting updrafts. There is no clear convergence of the results in terms of the mean tendency.

For the convective-scale contributions to the tendency, viz. $\rho u' w'$ averaged across the up/downrafts, the simulations do show convergence between the 1 km and 500 m grid spacing simulation (Fig. 9(b)). The 3 km case shows enhanced convective-scale fluxes above about 7 km altitude, consistent with less entrainment and larger undiluted updrafts (see e.g., (Bryan et al. 2003)). This apparent convergence at the convective scales, but not the domain-mean tendency suggests that at least some of the resolution sensitivity arises from changes in the mesoscale circulations with resolution.

These results compare favourably to those of Weisman et al. (1997), who found a similar behaviour for 2 km and 4 km grid spacing, which produced slightly greater momentum flux when compared to the 1 km simulation. Overall, the fluxes compared

well due to the coarser resolution model runs producing mature, upshear-tilted circulations. They imply that resolutions up to 12 km will produce mesoscale structures and therefore CMT need not be parametrised in such cases. However, those results are specifically focused on the mesoscale structure of squall lines and not the convective-scale structures as shown in Fig. 9(b). As discussed by Bryan et al. (2003), care should be taken not to use the model output from 1 km resolution simulations to represent a benchmark or control, as much finer resolution is needed, i.e. $\mathcal{O}(100 \text{ m})$, to produce turbulent flows including entrainment.

Another interesting point to note is the opposite sign of the momentum flux of the updrafts below 4 km and the downdrafts above 4 km. Figure 9(b) shows that the momentum flux from the downdrafts is positive above this height, and the updrafts are negative. Referring back to Fig. 2(c), the mesoscale momentum transport due to the front-to-rear flow and the rear-inflow jet (RIJ) have the same sign of $\rho u' w'$ (negative), which is consistent with the dominant updraft flux at mid- and upper levels and the downdraft flux at low-levels. The low-level countergradient updraft flux and upper-level countergradient downdraft flux likely arise from circulations that are tilted in the opposite direction to the convective system; Fig. 2(c) suggests that these circulations are located ahead of the convective system in the region of forward-tilted mesoscale circulations.

3.4. Effect of domain size on simulated convective systems

The sensitivity of the momentum transport to the size of the domain was investigated earlier to determine whether the assumptions underlying typical momentum transport parametrisation scheme are justified. Therefore, is it appropriate to approximate the tendency by the momentum flux only (as represented by Eqn. 6)? Also, is it appropriate to use small domains to study CMT as they struggle to represent the mesoscale transports, which we have shown in the previous section to be important? This is particularly relevant as it has become common to use long-running radiative-convective equilibrium (RCE) simulations with periodic boundary conditions to examine the structure and impact of moist convection, including momentum transport.

To explore this sensitivity, Hovmöller diagrams of cloud mixing ratio during the six hours simulation for all domain sizes with periodic boundary conditions were used to demonstrate how the different domains allow for the development of convection. In order to provide a direct comparison between the domains, for analysis purposes the domains were *trimmed* to the same size. The large and medium domains were trimmed to 100 km width and the square domain (50 km x 50 km) was duplicated. The cyclic simulations were also beneficial for this analysis because for open boundary conditions with small domains the convective system moves out of the domain. The largest two domains (Fig. 10 (top)) show the convective systems are maintained throughout the 6-hour simulation and beyond, and show very similar horizontal and temporal structure. Figure 10 (bottom) show that the small and the square domains have lifespans of around 2-3 hours only, dissipating thereafter. It is instructive to explore why these systems decayed earlier than those in the larger domains.

To determine what is causing these short-lived systems and why the convection in these small domains is consequently suppressed, the mean downward vertical velocity for the same (trimmed) regions shown in Fig. 10 was calculated. Figure 11 shows that the smaller domains have much stronger downward velocities, which suppresses the convection within those domains by stabilising the environment. The reason for this is that for periodic boundary conditions the domain-mean vertical velocity must be zero, and for a given convective system the strength of this subsidence must be larger for a smaller domain. As shown here for this

environment, once the domain gets too small (i.e. ≤ 100 km \times 50 km), the longevity of the convective system is reduced and it is unable to maintain a long-lived organised system.

One alternate explanation for the differences in organisation in the different sized domains is that the structure (and tilt) of the convective systems are different in their early stages, which cause them to fail to organise in the smaller domains. To determine whether the convective-scale behaviour of the convective systems in the smaller domains are unduly affected by the domain size at the early stages of their lifecycle, the momentum flux, the pressure gradient and the $\frac{1}{\rho} \frac{\partial \rho u'^2}{\partial x}$ terms averaged across the up/downdrafts for the small, medium and large domains were calculated during the first 2 hours of each simulation. As Fig. 12 indicates there is very little difference between the profiles from the various size domains. This implies that all the domains produce convective systems that are similar in structure during the initial stages of the simulation. Therefore in the first two hours all simulations have convective systems with similar convective-scale behaviour, and the small domains produce an adequate representation of the convective scales. It is the smaller domains, with their stronger compensating subsidence, that inhibits the transition of the convection to organised systems that can be maintained beyond a few hours.

Some recent studies (e.g. Romps (2012)) studied CMT using small domains approximately the size of our square domain and ran simulations to RCE. One conclusion of Romps's study was that the horizontal pressure gradient is not particularly important for the parametrisation of convective momentum transport. The results here show, however, that the small domain can produce systems with a notable cross-updraft pressure gradient (Fig. 12(a)), but these systems cannot be maintained. It is unlikely such a system would form spontaneously in a small-domain RCE simulation, which would instead be dominated by unorganised convection with weaker cross-updraft pressure gradients.

3.5. Evaluation of the Gregory parametrisation scheme

The GKI parametrisation scheme (Gregory et al. (1997)) is popularly used to parametrise CMT (with shallow and deep convection often parametrised separately (Stratton et al. (2009))). The GKI scheme was developed using results from Kershaw and Gregory (1997), which was used to estimate the momentum transports from various regimes of deep convection. Parametrised transports are determined using the mass-flux from the model's cumulus parametrisation scheme as well as a parametrisation of the effects of cross-updraft pressure gradients. A key aspect of this parametrisation is the assumption that the convective-scale pressure gradients (i.e., across the updrafts and downdrafts) are proportional to the product of the mass flux and the mean vertical wind shear, such that (for updrafts)

$$-\frac{\partial}{\partial x} \left(\frac{p'}{\rho} \right) \Big|_u \sim C_u M_u \frac{\partial \bar{u}}{\partial z} \quad (7)$$

where C_u is a constant, M_u is the mass flux across the updraft given by $\overline{\rho w' u'}$, p' is the pressure perturbation from the horizontal mean, \bar{u} represents the averages over the area covered by convective updrafts, and \bar{u} is the domain mean background wind. This relation was determined by linear theory and assumes that a high pressure anomaly exists on the upshear side on an updraft (see Rotunno and Klemp (1982)), akin to a plume in shear flow.

Figure 13 examines the terms in Eq. 7 across the updrafts, throughout the duration of the simulation using the medium domain with open boundary conditions. The mass flux M_u (Fig. 13a) is strictly positive (by definition) and accordingly the $M_u \frac{\partial \bar{u}}{\partial z}$ term (Fig. 13 (b)) is positive within the shear layer (< 4 km),

though varies in magnitude with time. Above the shear layer this term is small. The convective-scale pressure gradient (Fig. 13(c)) changes with both time and height, evolving from being negative at early times to being positive within the shear layer later, while remaining negative further aloft. This pressure variation, is broadly consistent with Eq. 7 in the first two hours, i.e. with the pressure gradient force being directed downshear, but as the system develops the pressure gradient force in the shear layer is directed upshear. This reversal of the low-level across-updraft pressure gradient is related to the development of the mesolow behind the leading convective line, as described by LeMone (1983), and is inconsistent with Eq. 7. This mesolow is also evident in Fig. 2(d), as indicated by the red contours.

The parameter C_u (Fig. 13(d)) (which is only shown here below 3.5 km as it is poorly defined further aloft) varies both in magnitude and sign. The time variation in the pressure gradient is mainly responsible for the fact that C_u is neither constant with height nor one-signed, contrary to an assigned fixed value. Moreover, C_u is only positive in the early stages of the system evolution before the system is organised. Grubišić and Moncrieff (2000) also found a variation of the value of C_u with height and though they found the GKI approximation to be accurate in the mid-levels, it lacked accuracy elsewhere. Grubišić and Moncrieff (2000) suggest that C_u should be a function of height z , as it varies between levels. The results presented here support this result, but further suggest that the GKI formulation is inadequate for organised systems.

Previously C_u , which is a tunable parameter, has been assigned various values depending on the model in which it is implemented. Values have ranged from 0.7 (Gregory et al. (1997)) which is also used by UKMO (Stratton et al. (2009)) and an earlier version of NCAR's Community Atmosphere Model (CAM) 3.0 (Richter and Rasch (2008)); 0.55 (Zhang and Wu (2003)); and recently, 0.4 in CAM 5.1 (Neale et al. (2010)). Over the 6-hour model run shown in Fig. 13(d) the average value of C_u equals approximately -0.3 within the shear layer, which is smaller (in magnitude) and of the opposite sign to all of the above values. In the first two hours of the simulation, however, C_u is approximately 0.5, which is consistent with the above values.

As noted by Neale et al. (2010), the magnitude of C_u (and correspondingly C_d) has an important control on the strength of parametrised convective momentum transport as the pressure gradient is a 'sink' term; as C increases the strength of the parametrised momentum transport decreases. However, as shown earlier (e.g., Fig. 5) the pressure gradient can actually act in concert with the up/down draft fluxes, such as at low-levels in the early stages of the evolution of the convective system. Ultimately, this pressure gradient evolves with the convective system and as shown for this organised system is poorly represented by Eq. 7. This poor representation is because Eq. 7 does not take into account the upshear directed pressure gradient associated with the elevated mesolow behind the leading edge of organised systems.

3.6. Evaluation of the Moncrieff mesoscale momentum transport models

The Moncrieff analytic models of organised moist convection are based on the conservation of mass, entropy, total energy, vorticity and horizontal momentum generation for inviscid steady flow in Lagrangian space (Moncrieff (1981)). These nonlinear models are approximately exact because the sole assumption is that the buoyancy is a separable function of the vertical velocity, which is valid for moist adiabatic motion. These models have been comprehensively evaluated by cloud-system resolving model simulation data and field-campaign measurements (e.g. Houze

(2004), Houze (2014), Moncrieff (2010)). In particular, the two-dimensional analytic models are solutions of the elliptic integro-differential vorticity equation:

$$\nabla^2 \psi = G(\psi) + \int_{z_0}^z \left(\frac{\partial F}{\partial \psi} \right)_{z'} dz' \quad (8)$$

where ψ is the streamfunction, G is the environmental shear and the integral is the vorticity generated by the horizontal gradient of buoyancy.

The far-field solution of Eq. 8 provides open lateral boundary conditions for the two-dimensional models. Analytic models of three-dimensional propagating squall lines (Moncrieff and Miller (1976)), and upshear/downshear propagating convective bands (Lane and Moncrieff (2015)/Moncrieff and Lane (2015)) are tractable only in the far-field. However, the Lagrangian formulation of the analytic models (see below for the two-dimensional models) enables the three-dimensional transports to be calculated from the far-field solutions.

The two-dimensional Moncrieff (1992) archetypal model (Fig. 14), the minimalist mesoscale system, is governed by the quotient of work done by the horizontal pressure gradient and the inflow kinetic energy, $E = \Delta p_s / \frac{1}{2} \rho U_s^2$, where Δp_s and U_s are the surface pressure change across the system and the surface inflow speed, respectively i.e., the ratio of the two terms of the Bernoulli pressure. With application to MCSs, squall lines, and density currents the archetypal models have a distinguished rearward-tilted circulation, and the mesoscale momentum transport has sign opposite to that of the propagation speed, e.g., momentum transport is negative when the propagation speed is positive, and vice versa.

The Lagrangian based two-dimensional analytic models enables the momentum transport to be derived from the far-field solutions of the vorticity equation. The two-dimensional archetypes is as follows. Referring to Fig. 14, the momentum transport is for the total relative flow (u_m, w_m) . Define the difference operator $\Delta = [\]_0^{L_m}$ and the horizontal averaging operator $\langle \rangle = \frac{1}{L_m} \int_0^{L_m} (\) dx$. Integration of the steady Eulerian momentum equation in the x-direction

$$\frac{\partial}{\partial x} \left(u_m^2 + \frac{p_m}{\rho} \right) + \frac{\partial}{\partial z} (u_m w_m) = 0 \quad (9)$$

gives

$$\frac{\partial}{\partial z} \langle u_m w_m \rangle = -\frac{1}{L_m} \Delta \left[u_m^2 + \frac{p_m}{\rho} \right]_0^{L_m} \quad (10)$$

In the above, u_m and w_m are normalised by U_s , and $\frac{p_m}{\rho}$ by U_s^2 . The mesoscale momentum transport divergence is proportional to the cross-system change of Bernoulli pressure and the aspect ratio ($\frac{1}{L_m}$) is the constant of proportionality. Integration of Eq. 10 with $w_m = 0$ at the horizontal lower boundary gives the mesoscale momentum transport at height z .

$$\langle u_m w_m \rangle = -\frac{1}{L_m} \int_0^z \Delta \left[u_m^2 + \frac{p_m}{\rho} \right]_0^{z'} dz' \quad (11)$$

The $w_m = 0$ boundary condition at the horizontal upper boundary provides a powerful integral constraint on the mesoscale momentum transport. It follows that the Bernoulli pressure

constrains the momentum tendency

$$\int_0^1 \Delta \left[u_m^2 + \frac{p_m}{\rho} \right]_0^{L_m} dz = 0 \quad (12)$$

i.e., the horizontal momentum is redistributed but the net-momentum generation is zero.

There are important distinctions between the mesoscale momentum transport based on the analytic models and the traditional eddy-based momentum transport, where an eddy is defined as a deviation from the horizontal mean (see Eq. 4). In other words, the eddy flux is replaced by the total mesoscale transport divergence, i.e., eddy + mean components. Therefore, the mesoscale acceleration of the mean flow vertical profile is

$$\frac{\delta}{\delta t} \langle u_m \rangle = -\alpha_m \frac{\partial}{\partial z} \langle u_m w_m \rangle = \alpha_m \frac{1}{L_m} \left[u_m^2 + \frac{p_m}{\rho} \right]_0^{L_m} \quad (13)$$

where α_m is an amplitude function or closure (Moncrieff (1992), Eq. 24). Fig. 13, showed that the momentum transport divergence for the archetypal model accelerates/decelerates the mean flow in the lower/upper troposphere, respectively. The above formulae can be evaluated using the numerical simulations herein with open lateral boundary conditions. The terms in Eqn. 10 are represented in Fig. 15 for the medium open domain with 1 km grid spacing.

An important structural characteristic of the archetypal models is that the upshear tilt of the mesoscale circulation means that the sign of the mesoscale transport is opposite to the propagation vector, e.g., a system propagating in the positive x-direction has negative momentum transport. The archetypal model is controlled by $E = \Delta p / \frac{1}{2} \rho U_s^2$, the ratio of the two components of Bernoulli pressure at the lower boundary. Figure 2 of Moncrieff (1992) shows the morphology of the relative flow for systems in the range $-8 \leq E \leq \frac{8}{9}$. The lower and upper limits represent a system with no downdraft and no overturning updraft (strict propagation), respectively. Systems featuring deep downdrafts and deep overturning updrafts are associated with small absolute values of E . This structure pertains to the numerical simulations because Δp is small and $U_s = -18 \text{ ms}^{-1}$ so E is approximately zero. Vertical integration of Eq. 17 in Moncrieff (1992) gives the momentum transport for $E = 0$ as

$$\langle u_m w_m \rangle = \frac{4}{3} \left(\frac{z}{h} \right)^3 - 2 \left(\frac{z}{h} \right)^2 \quad (14)$$

for $0 \leq z \leq h$, and provides values for $h \leq z \leq 1$ due to the symmetry of the profile. The archetypal model momentum transport profile shown in Fig. 16(a) resembles the transport calculated from the numerical simulation in Fig. 16(b). In dimensional units the minimum value of $\langle u_m, w_m \rangle$ in the archetypal model is $-13.6 \text{ m}^2 \text{ s}^{-2}$, using $L_m = 200/12 = 16.7$ and $U_s = -18 \text{ ms}^{-1}$. This is in good agreement with $-15.5 \text{ m}^2 \text{ s}^{-2}$ for the simulation results (Fig. 16(b)).

4. Synthesis and Conclusion

This study on the momentum budget of idealised convective systems focused on aspects relevant for the parametrisation of convective momentum transport. Specifically, there are important sensitivities of the simulated momentum budget to model domain size and resolution, as well as the contributions from convective-scale transports to the overall budgets.

Simulations with varied domain size revealed the key role of horizontal gradient terms in contributing to the momentum

1 budget. In particular, even for relatively large domains (viz. the
2 'medium' domain that was 400 km long) the horizontal pressure
3 gradient was a notable contributor to the mean flow tendency. This
4 mesoscale pressure gradient arose entirely from the simulated
5 convective system, acting to maintain the organised circulation
6 associated with that system. This mesoscale pressure gradient
7 mostly opposes the tendency from the convective momentum
8 flux divergence, and for the 'medium' domain was about half
9 the magnitude of the momentum flux tendency. Even for the
10 'large' domain (length = 800 km), the mesoscale pressure gradient
11 was about one quarter the magnitude of the momentum flux
12 term. This specific result has important implications for CMT
13 parametrisation as these domain sizes are larger than most
14 global model grid boxes; this mesoscale pressure gradient is not
15 represented in any scheme, but it clearly plays an important role
16 for mesoscale systems. Even when the work done by the static
17 pressure ($\frac{p}{\rho}$) term is identically zero; namely $E=0$, the system still
18 has a distinguished vertical tilt (see Moncrieff (1992); Fig. 4) that
19 provides mesoscale momentum transport.

20 The 'small' domain simulations considered here were not
21 large enough for the convective systems to evolve and maintain
22 themselves realistically. As shown in Fig. 12, the systems across
23 all domains develop in a similar fashion for the first few hours.
24 The pressure gradient and the structure of the convective system
25 in these small domains are similar during the first two hours of
26 the simulation (Fig. 5 and Fig. 10). However, when the domains
27 are too small ($O(\leq 100 \text{ kms})$), then the convection is suppressed
28 by overly strong subsidence and is unable to become properly
29 organised. This has important implications for the interpretation
30 of momentum fluxes associated with convection in small-domain
31 RCE simulations e.g., Romps (2012), as they are unable to
32 properly represent organisation.

33 The effect of model resolution on the CMT showed the
34 3 km grid spacing model producing larger fluxes and more
35 convective overshoots than higher resolution domains. This result
36 was consistent with previous studies that attribute some of
37 the sensitivities to insufficient entrainment at these convection-
38 permitting resolutions. Though, at least for those simulations
39 presented here, there seemed to be convergence of the convective-
40 scale fluxes at grid spacings of 1 km.

41 Comparison of the convective-scale transports, i.e., those
42 associated with individual updrafts and downdrafts, and those
43 associated with the domain mean finds that in the early stages of
44 evolution the convective-scales work against the mesoscales and
45 result in a near-zero tendency. At mature stages, after the system
46 tilts upshear, the convective-scale and mesoscale transports act in
47 concert in a downgradient manner. As part of this evolution the
48 low-level cross-draft pressure gradients change sign during the
49 system evolution associated with the development of organised
50 mesoscale circulations.

51 The most common parametrisations of convective momentum
52 transport, which are based on the GKI scheme, use entraining
53 plume models and incorporate a simple representation of the
54 effects of the cross-updraft pressure gradient on the momentum
55 tendency. As shown here, this representation poorly reproduces
56 the evolution of the pressure gradient, because the assumed
57 constant of proportionality between the pressure gradient and the
58 product of the mean shear and the mass-flux actually changes
59 sign as the convective system matures. This was demonstrated
60 by comparing the convective-scale transports associated with
individual updrafts and downdrafts with the domain mean (e.g.
Fig. 4). In the early stages of the convective evolution, the
negative convective-scale momentum flux is offset by the positive
mesoscale momentum flux which leads to a near-zero total
tendency. After the mature stage when the system develops an
upshear tilt, the sign of the mesoscale momentum flux changes

from positive to negative meaning the two momentum transports
act in concert, consistent with the change in sign of the low-level
horizontal pressure gradient.

The importance of introducing organised mesoscale transport
to CMT has recently been underlined by global climate
model experiments. Moncrieff et al. (2017) implemented
dynamically based parametrisations of convective heating and
mesoscale momentum in the Community Atmosphere Model.
Both parametrisations significantly impacted, in distinct ways,
convectively coupled equatorial waves and the large-scale
distribution of tropical precipitation in the warm-pool and the
adjoining regions that commonly feature multiscale convective
organisation. From those results and the results presented herein it
is clear that the mesoscale contributions to the net momentum flux
are complicated, and can offset the contributions from convective
scales. Thus, an approach that goes beyond the standard plume
model is required to represent the momentum tendencies from
mesoscale convection systems properly, and should be a priority
for future parameterisation efforts.

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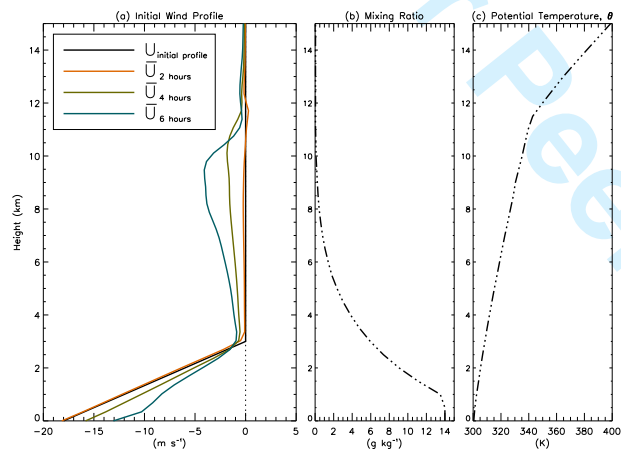


Figure 1. Profiles of (a) zonal wind u (m s^{-1}), (b) mixing ratio q (g kg^{-1}) and (c) potential temperature θ (K) used in the initial sounding for all simulations. Also shown in (a) is the mean wind \bar{u} every two hours for the medium 1 km domain with cyclic boundary conditions.

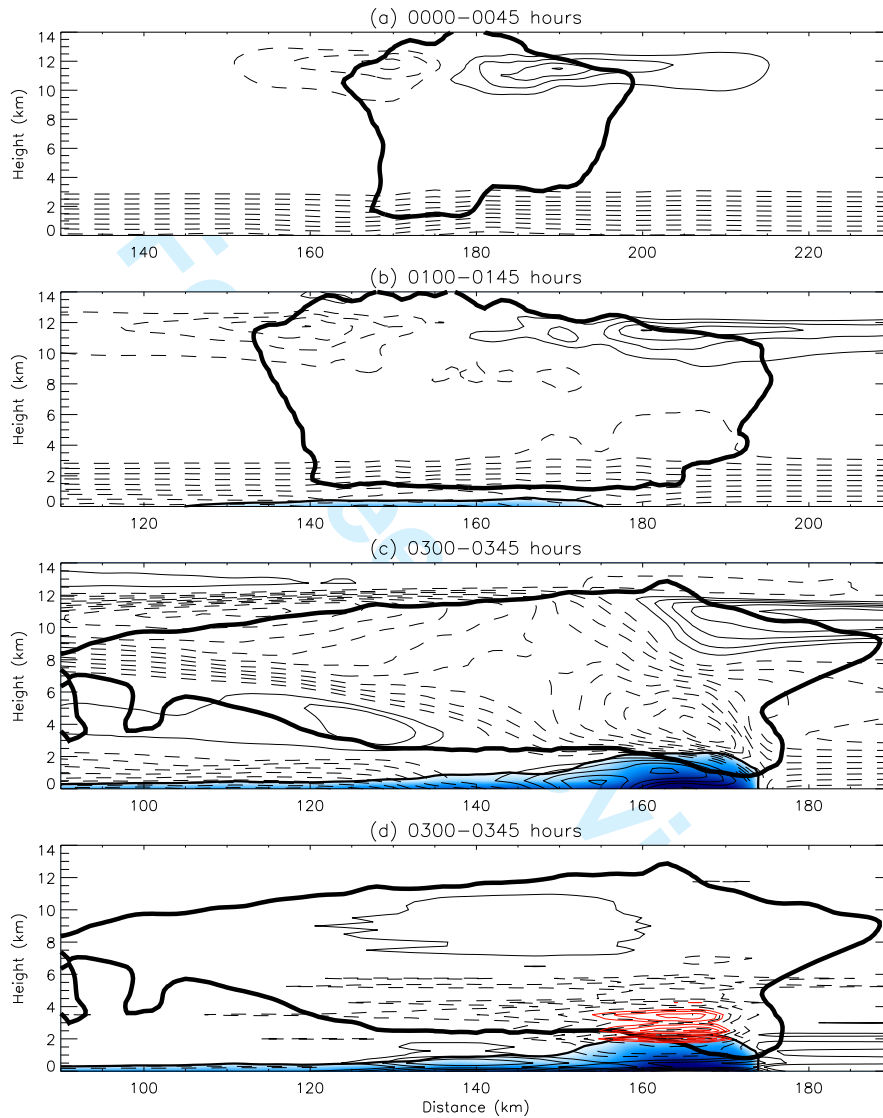


Figure 2. Cross-section of cloud (cloud mixing ratio contour outline 0.1 g kg^{-1} is shown by solid black line) averaged in the y -direction and time-averaged over the following time periods: (a) 0000-0045 hours, (b) 0100-0145 hours and (c) & (d) 0300-0345 hours when the system is fully developed. (a)-(c) shows contours of u (every 2.0 m s^{-1} , with zero removed) - positive values shown by solid lines and negative by dashed) for medium 1 km resolution cyclic domain. The cold pool is shown by blue shading, using an outer contour of -1°C temperature perturbation. (d) The black contours are pressure perturbations every 30 Pa , with negative values shown by dashed lines and positive values shown by solid lines. The red contours shows the pressure perturbations every 30 Pa , less than 150 Pa i.e. the region of lowest pressure.

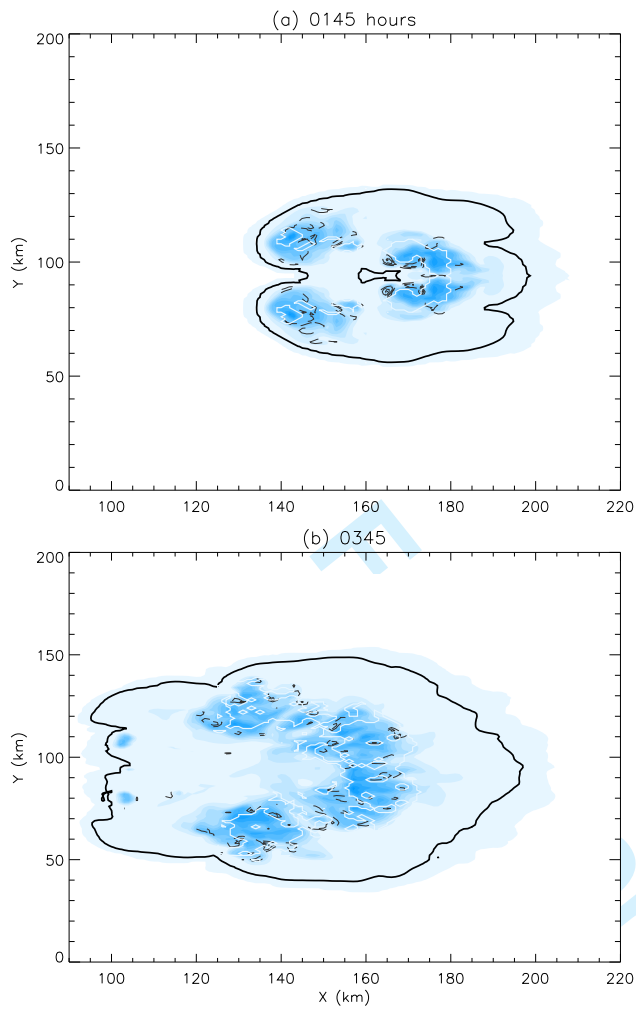


Figure 3. Horizontal cross-section of cloud at 8 km height at (a) 0145 hours and (b) 0300 hours. Cloud outline is cloud mixing ratio 0.1 g kg^{-1} . Black contours represent areas of convective updrafts (defined as in-cloud $w > 4 \text{ m s}^{-1}$ and 15-min surface rain $> 1 \text{ mm h}^{-1}$). The blue shading represents the cloud mixing ratio, the darkest shades represent values of 1.1 g kg^{-1} and the lightest shading is 0.05 g kg^{-1} .

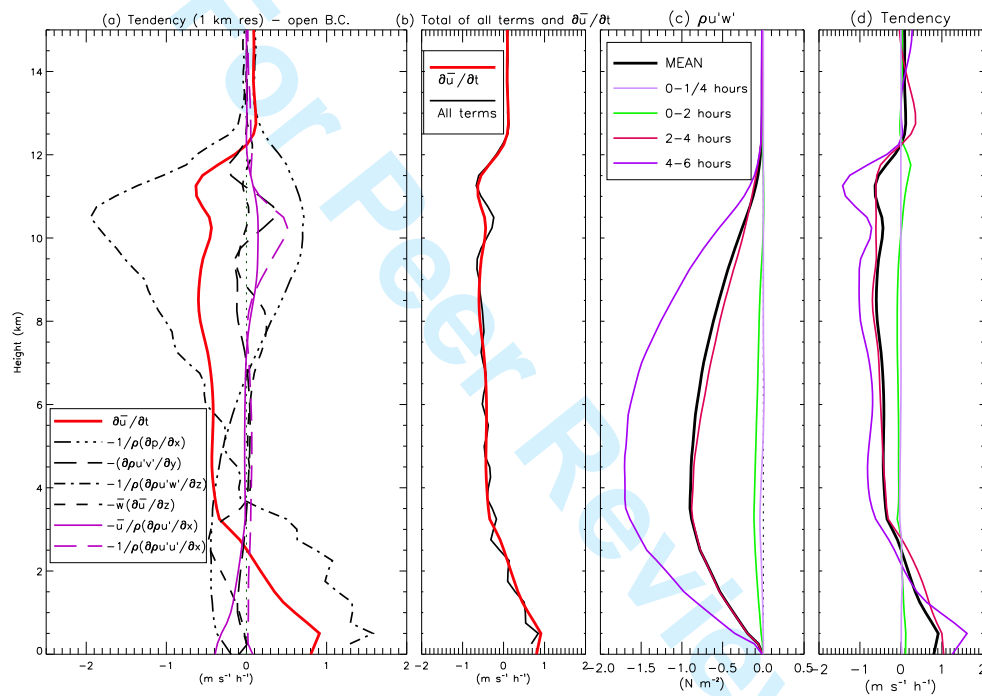


Figure 4. Profiles of (a) the domain mean momentum tendency (red line) and all contributing terms ($\text{m s}^{-1} \text{h}^{-1}$) averaged over the 6 h simulation, (b) comparison of the momentum tendency (red line) and the sum of all the terms (black line) ($\text{m s}^{-1} \text{h}^{-1}$), (c) momentum transport $\rho u'w'$ (N m^{-2}) and (d) the tendency in $\text{m s}^{-1} \text{h}^{-1}$ for the 1 km horizontal resolution, medium domain with open boundary conditions. (c) and (d) show various stages throughout the 6 h simulation.

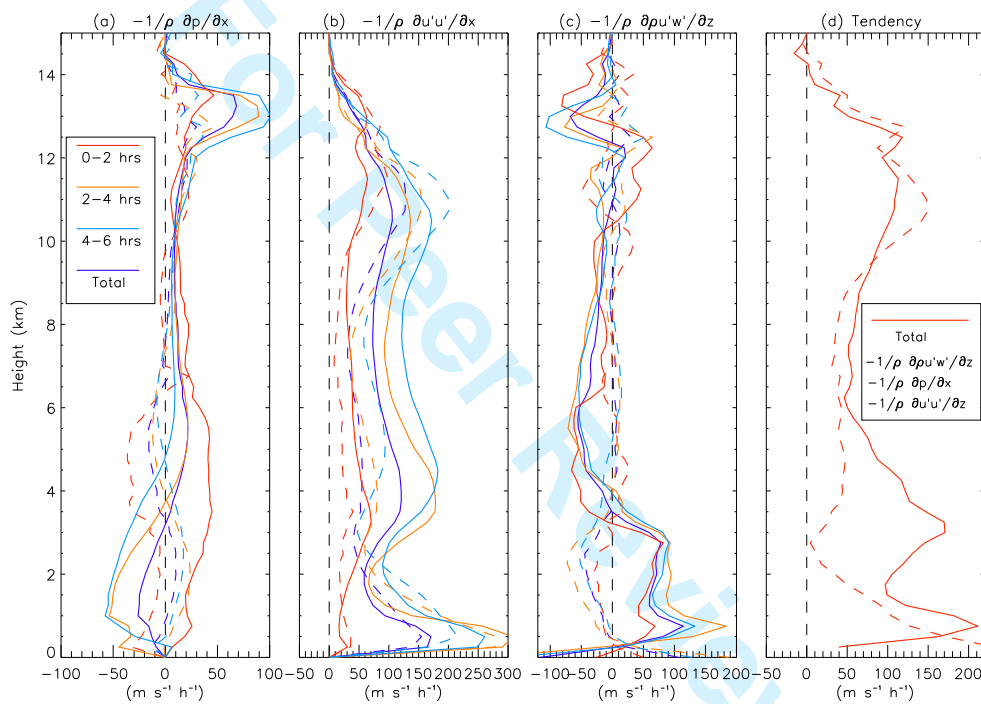


Figure 5. Time-averaged terms that have been area-averaged across the updrafts and downdrafts, defined as in-cloud and $w > 1 \text{ m s}^{-1}$ and $< -1 \text{ m s}^{-1}$, respectively. (a) pressure gradients $-\frac{1}{\rho} \frac{\partial p}{\partial x}$, (b) $-\frac{1}{\rho} \frac{\partial \rho u' u'}{\partial x}$, (c) $-\frac{1}{\rho} \frac{\partial \rho u' w'}{\partial z}$ and (d) the sum of $-\frac{1}{\rho} \frac{\partial \rho u' w'}{\partial z} - \frac{1}{\rho} \frac{\partial p}{\partial x} - \frac{1}{\rho} \frac{\partial \rho u' u'}{\partial x}$ which represents the time-varying tendency, for the 1 km horizontal resolution, medium domains with open boundary conditions. Dashed lines are the downdraft profiles and the solid lines are the updraft profiles.

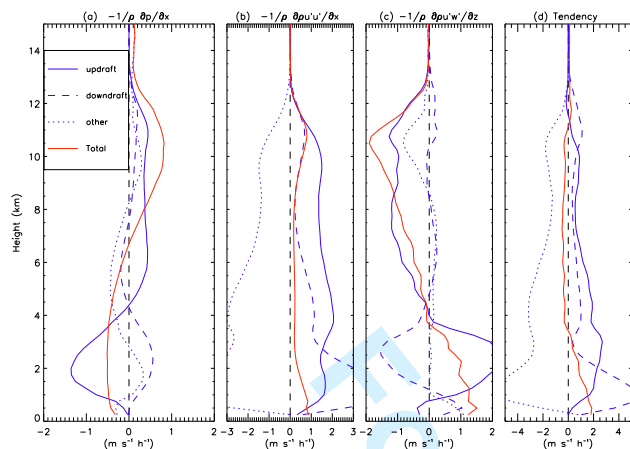


Figure 6. Relative contributions of terms shown in Fig. 5 to the total profile of (a) pressure gradients $-\frac{1}{\rho} \frac{\partial p}{\partial x}$, (b) $-\frac{1}{\rho} \frac{\partial \rho u' u'}{\partial x}$, (c) $-\frac{1}{\rho} \frac{\partial \rho u' w'}{\partial z}$ and (d) the sum of $-\frac{1}{\rho} \frac{\partial \rho u' w'}{\partial z} - \frac{1}{\rho} \frac{\partial p}{\partial x} - \frac{1}{\rho} \frac{\partial \rho u' u'}{\partial x}$ which represents the time-varying tendency.

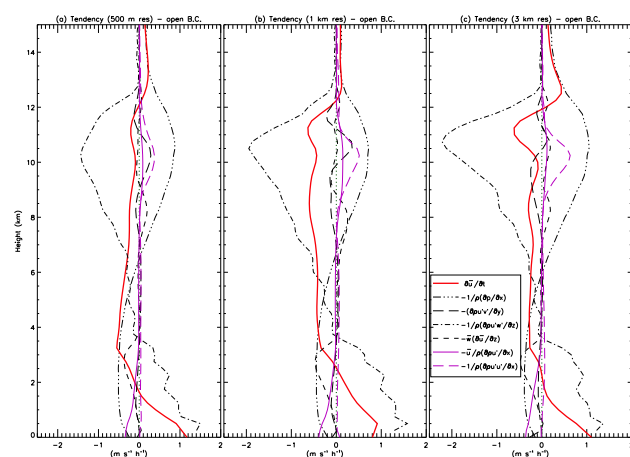


Figure 7. Momentum budgets for 1 km medium (top) and large (below) domains with open (left) and cyclic (right) boundary conditions. The total tendency $\frac{\partial \bar{u}}{\partial t}$ is shown in red.

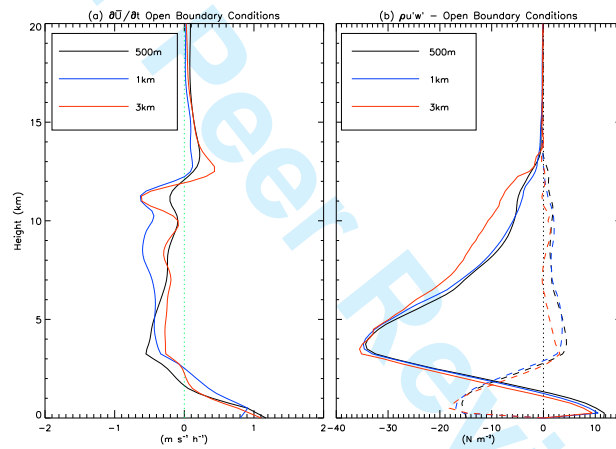


Figure 8. Momentum budget comparison for 500 m (left), 1 km (centre) and 3 km (right) models, for medium domain with open boundaries. Domain average values of terms shown in key.

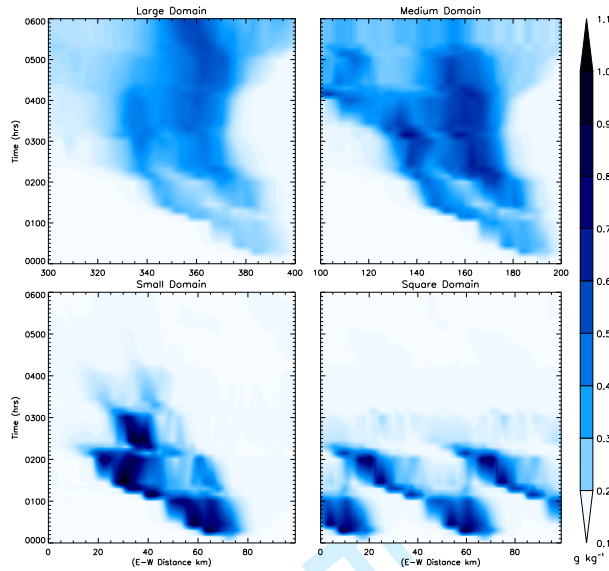


Figure 9. Comparison of the tendency (left) and momentum (right) for open boundary conditions for the various resolution medium domains. The momentum transport contributions of updrafts (solid line) and downdrafts (dashed line) are shown.

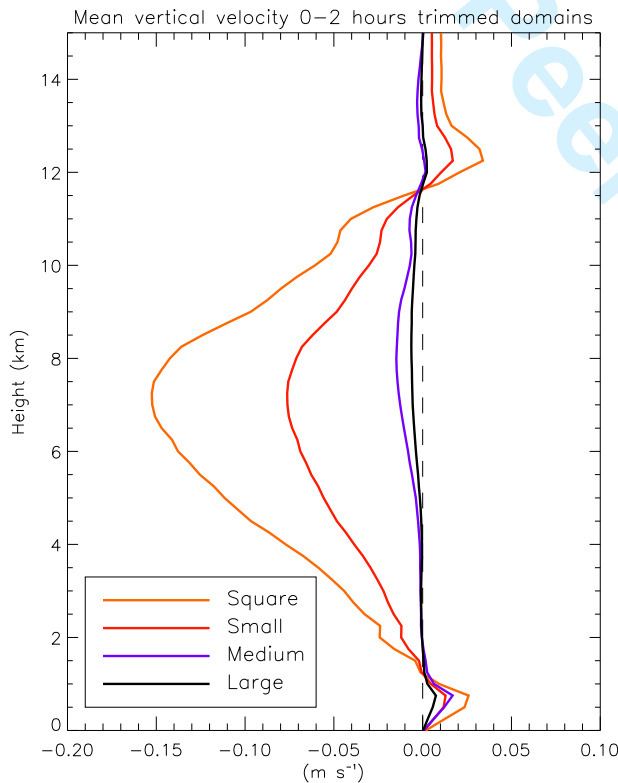


Figure 10. Hovmöller diagrams of cloud for the 1 km cyclic, large and medium domains (top left and right) and small and square domains (below left and right). The square domain was doubled to have the same 100 km width as the other domains (the large and medium domains are trimmed to 100 km). The blue shading shows the average mixing ratio (cloud, rain, ice, snow and graupel) in g kg^{-1} .

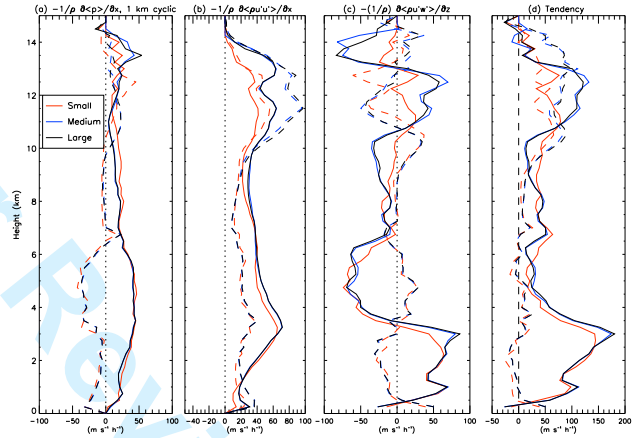


Figure 11. Mean vertical velocity out of cloud across trimmed domains (100 km width) for all cyclic domains at 1 km resolution.

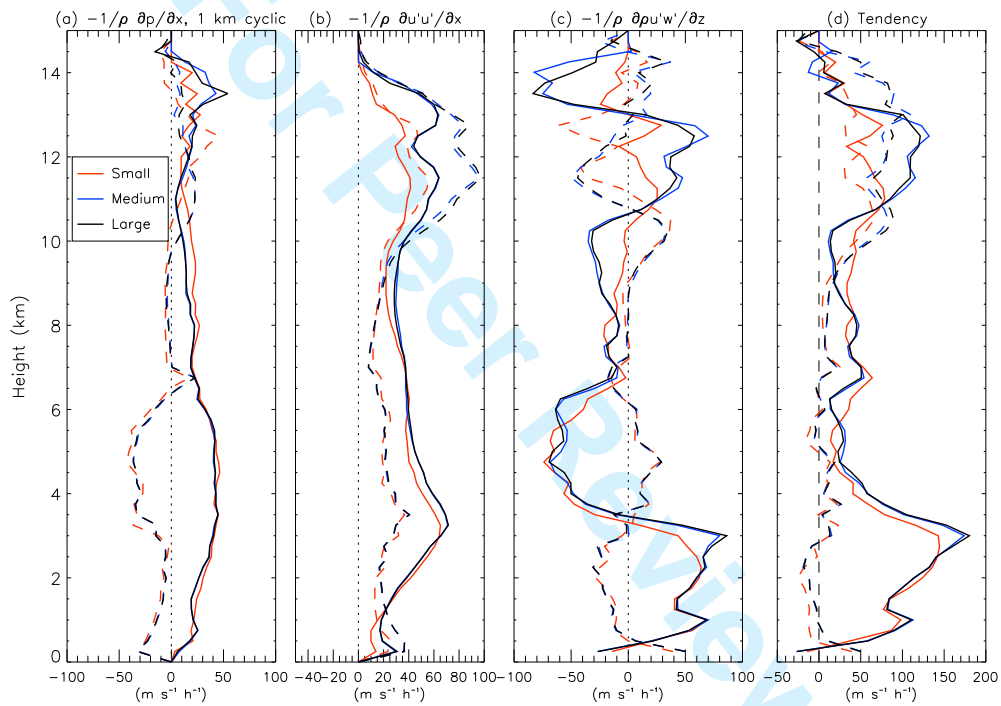


Figure 12. Time-averaged (a) pressure gradients $-\frac{1}{\rho} \frac{\partial p}{\partial x}$, (b) $-\frac{1}{\rho} \frac{\partial \rho u' u'}{\partial x}$, (c) $-\frac{1}{\rho} \frac{\partial \rho u' w'}{\partial z}$ and (d) $-\frac{1}{\rho} \frac{\partial \rho u' w'}{\partial z} - \frac{1}{\rho} \frac{\partial p}{\partial x} - \frac{1}{\rho} \frac{\partial \rho u' u'}{\partial x}$. All terms are averaged across updrafts and downdrafts for the 1 km horizontal resolution domains with cyclic boundary conditions during 0–2 hours.

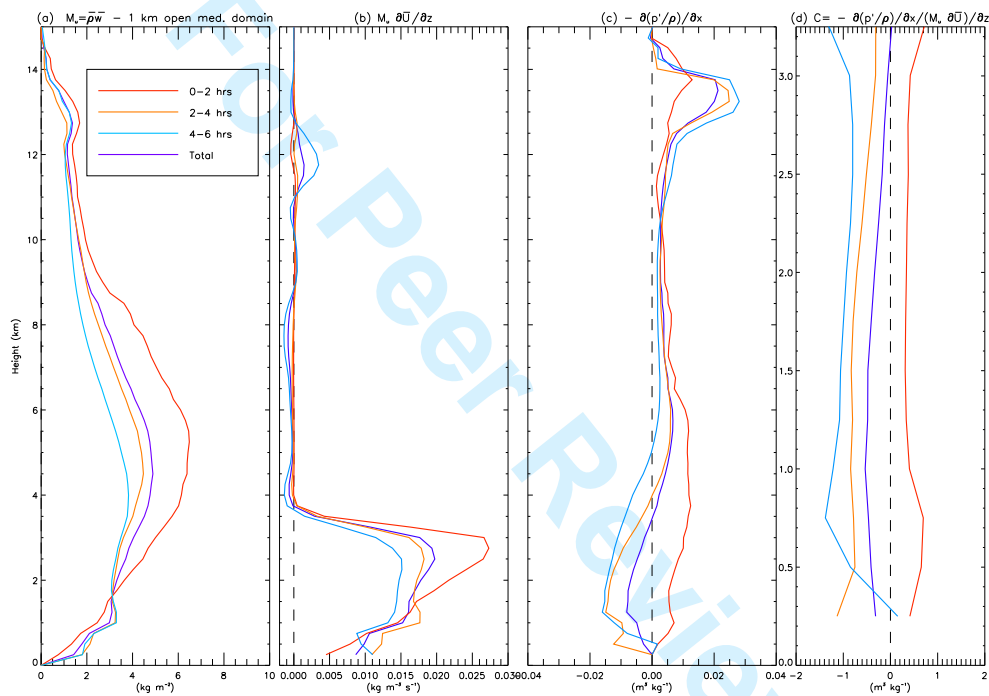


Figure 13. Terms in the relation used in the Gregory parametrisation scheme, $-\frac{\partial}{\partial x} \left(\frac{p'}{\rho} \right) \propto C_u M_u \frac{\partial \bar{U}}{\partial z}$ - shows (a) $M_u = \overline{\rho w}$, (b) $M_u \frac{\partial \bar{U}}{\partial z}$, (c) $-\frac{\partial}{\partial x} \frac{p'}{\rho}$ and (d) C_u where $C_u = \frac{-\frac{\partial}{\partial x} \frac{p'}{\rho}}{M_u \frac{\partial \bar{U}}{\partial z}}$ but only up to 3.25 km - (the height of the shear level) for the 1 km resolution medium open domain. These terms are averaged across time periods and across updrafts ($> 1 \text{ m s}^{-1}$) for each height level.

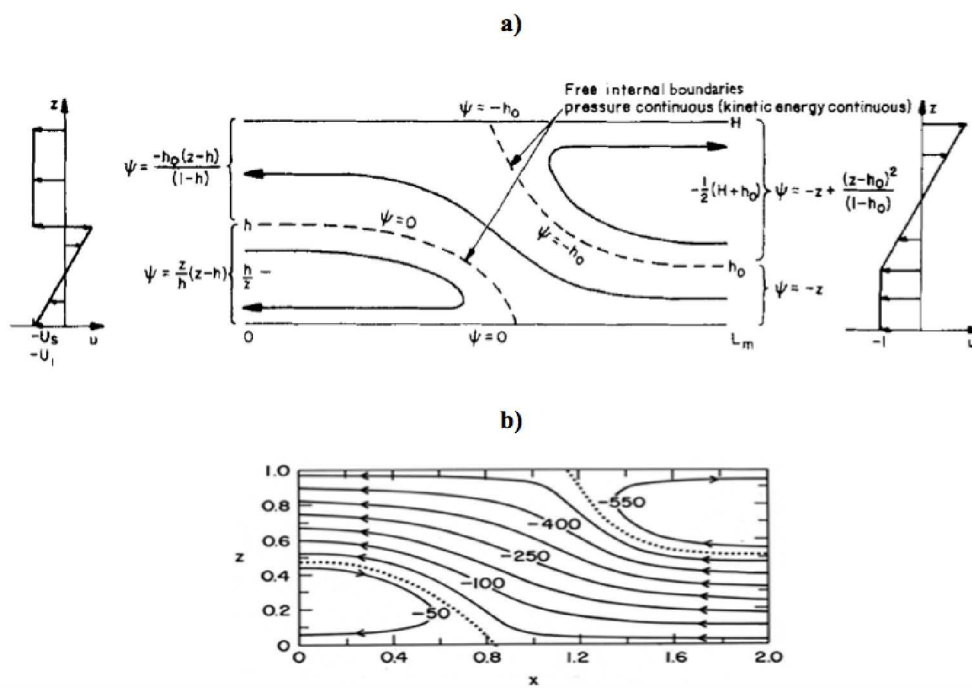


Figure 14. a) Schematic description of the Moncrieff (1992) archetypal model based on the conservation of mass, total energy, vorticity in a Lagrangian framework, and the integral constraint on momentum transport. b) Numerical solution of the two-dimensional vorticity equation shows the characteristic rearward-tilt of the airflow. The dotted and broken lines denote free-boundaries whose shapes are part of the two-dimensional solution, subject to boundary conditions defined by the far-field solution. Fig. a) is from Moncrieff (1992).

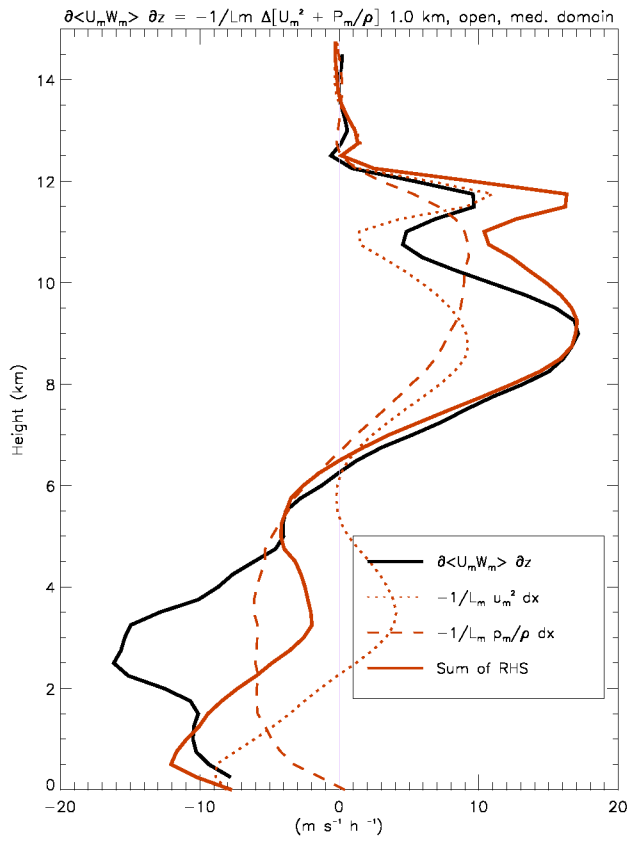


Figure 15. Terms in the relationship $\frac{\partial}{\partial z} \langle u_m w_m \rangle = -\frac{1}{L_m} \Delta \left[u_m^2 + \frac{p_m}{\rho} \right]_{0}^{L_m}$ for the open lateral boundary conditions for medium domain. Averaged over 0400-0500 hours during the mature stage of evolution. Sum of the terms on the RHS are shown by the solid red line.

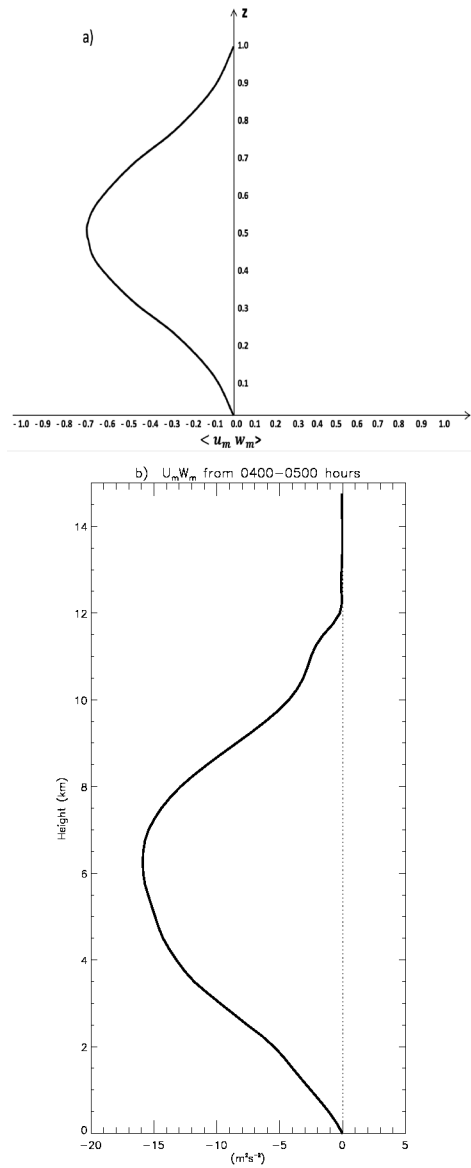


Figure 16. Top - a) Profile of $\langle u_m w_m \rangle$ for archetypal model in Moncrieff (1992). Below - b) Profile of $u_m w_m$ from the numerical simulation using the medium open domain.



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